

# Insights on Atmospheric Oxidation Processes by Performing Factor Analyses on Sub-ranges of Mass Spectra

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## Abstract

Our understanding of atmospheric oxidation chemistry has improved significantly in recent years, greatly facilitated by developments in mass spectrometry. The generated mass spectra typically contain vast amounts of information on atmospheric sources and processes, but the identification and quantification of these is hampered by the wealth of data to analyze. The implementation of factor analysis techniques have greatly facilitated this analysis, yet many atmospheric processes still remain poorly understood. Here, we present new insights on highly oxygenated products from monoterpene oxidation, measured by chemical ionization mass spectrometry, at a boreal forest site in Finland in fall 2016. Our primary focus was on the formation of accretion products, i.e. “dimers”. We identified the formation of daytime dimers, with a diurnal peak at noon time, despite high nitric oxide (NO) concentrations typically expected to inhibit dimer formation. These dimers may play an important role in new particle formation events that are often observed in the forest. In addition, dimers identified as combined products of NO<sub>3</sub> and O<sub>3</sub> oxidation of monoterpenes were also found to be a large source of low-volatile vapors at night. This highlights the complexity of atmospheric oxidation chemistry, and the need for future laboratory studies on multi-oxidant systems. Neither of these two processes could have been separated without the new analysis approach deployed in our study, where we applied binned positive matrix factorization (binPMF) on sub-ranges of the mass spectra, rather than the traditional approach where the entire mass spectrum is included for PMF analysis. In addition to the main findings listed above, several other benefits compared to traditional methods were found.

## 34 **1 Introduction**

35 Huge amounts of volatile organic compounds (VOC) are emitted to the atmosphere every year  
36 (Guenther et al., 1995;Lamarque et al., 2010), which play a significant role in atmospheric chemistry  
37 and affect the oxidative ability of the atmosphere. The oxidation products of VOC can contribute to  
38 the formation and growth of secondary organic aerosols (Kulmala et al., 2013;Ehn et al., 2014;Kirkby  
39 et al., 2016;Troestl et al., 2016), affecting air quality, human health, and climate radiative forcing  
40 (Pope III et al., 2009;Stocker et al., 2013;Zhang et al., 2016;Shiraiwa et al., 2017). Thanks to the  
41 advancement in mass spectrometric applications, like the aerosol mass spectrometer (AMS)  
42 (Canagaratna et al., 2007) and chemical ionization mass spectrometry (CIMS) (Bertram et al.,  
43 2011;Jokinen et al., 2012;Lee et al., 2014), our capability to detect these oxidized products, as well  
44 as our understanding of the complicated atmospheric oxidation pathways in which they take part,  
45 have been greatly enhanced.

46 Monoterpenes ( $C_{10}H_{16}$ ), one major group of VOC emitted in forested areas, have been shown to be a  
47 large source of atmospheric secondary organic aerosol (SOA). The oxidation of monoterpenes  
48 produces an abundance of different oxidation products (Oxygenated VOC, OVOC), including highly  
49 oxygenated organic molecules (HOM) with molar yields in the range of a few percent, depending on  
50 the specific monoterpene and oxidant (Ehn et al., 2014;Bianchi et al., 2019). Recent chamber studies  
51 have greatly advanced our knowledge of formation pathways for monoterpene HOM products, e.g.  
52 “monomers” (typically  $C_{9-10}H_{12-16}O_{6-12}$ ) and “dimers” (typically  $C_{19-20}H_{28-32}O_{8-18}$ ). Dimers, as shown  
53 by previous studies, can contribute to new particle formation (NPF) (Kirkby et al., 2016;Troestl et al.,  
54 2016;Lehtipalo et al., 2018), and are thus of particular interest.

55 In nearly all atmospheric oxidation chemistry, peroxy radicals ( $RO_2$ ) are the key intermediates  
56 (Orlando and Tyndall, 2012). They form when VOC react with oxidants like ozone, or the hydroxyl  
57 (OH) or nitrate ( $NO_3$ ) radicals, while their termination occurs mainly by bimolecular reactions with  
58 nitric oxide (NO), hydroperoxyl ( $HO_2$ ) and/or other  $RO_2$ .  $RO_2+R'O_2$  reactions can form  $ROOR'$   
59 dimers (Berndt et al., 2018a;Berndt et al., 2018b), and this pathway competes with  $RO_2+NO$  reactions,  
60 meaning that NO, formed by photolysis of  $NO_2$ , can efficiently suppress dimer formation, as also  
61 seen from atmospheric HOM observations (Ehn et al., 2014;Yan et al., 2016). Mohr et al. (2017) also  
62 reported daytime dimers in the boreal forest in Finland, coinciding with NPF events. A better  
63 understanding of the formation of these daytime dimers would assist elucidating NPF and particle  
64 growth mechanisms.

65 At night, nitrogen oxides can also impact the oxidation pathways, when  $NO_2$  and  $O_3$  react to form  
66  $NO_3$  radicals that can oxidize monoterpenes.  $NO_3$  radicals are greatly reduced during daytime due to  
67 photolysis and reactions with NO reducing their lifetime to a few seconds (Ng et al., 2017). Yan et

68 al. (2016) reported nighttime HOM initiated by NO<sub>3</sub> in the boreal forest in Finland, but to our  
69 knowledge there have been no laboratory studies on HOM formation from NO<sub>3</sub> oxidation of  
70 monoterpenes. However, there have been several studies looking into the SOA formation in these  
71 systems, finding that certain monoterpenes, like β-pinene, have very high SOA yields, while the most  
72 abundant monoterpene, α-pinene, has negligible SOA forming potential. It remains an open question  
73 what the role of NO<sub>3</sub> radical oxidation of monoterpenes, and the observed NO<sub>3</sub>-derived HOM, in the  
74 night-time boreal forest is. Identification of these processes in the ambient environment is  
75 fundamental towards better understanding of NPF and SOA.

76 The recent development of CIMS techniques has allowed researchers to observe unprecedented  
77 numbers of OVOC, in real-time (Riva et al., 2019). This ability to measure thousands of compounds  
78 is a great benefit, but also a large challenge for the data analyst. For this reason, factor analytical  
79 techniques have often been applied to reduce the complexity of the data (Huang et al., 1999), e.g.  
80 positive matrix factorization, PMF (Paatero and Tapper, 1994; Zhang et al., 2011). The factors have  
81 then been attributed to sources (e.g. biomass burning organic aerosol) or processes (e.g. monoterpene  
82 ozonolysis) depending on the application and ability to identify spectral signatures (Yan et al.,  
83 2016; Zhang et al., 2017).

84 In the vast majority of these PMF applications to mass spectra, the mass range of ions has been  
85 maximized in order to provide as much input as possible for the algorithm. This approach was  
86 certainly motivated in early application of PMF on e.g. offline filters, with chemical information of  
87 metals, water-soluble ions, and organic and elemental carbon (OC and EC), where the number of  
88 variables is counted in tens, and the number of samples in tens or hundreds (Zhang et al., 2017).  
89 However, with gas-phase CIMS, we often have up to a thousand variables, with hundreds or even  
90 thousands of samples, meaning that the amount of data itself is unlikely to be a limitation for PMF  
91 calculation. In this work, we aimed to explore potential benefits of dividing the spectra into sub-  
92 ranges before applying factorization analysis. This approach was motivated by several issues, which  
93 we expected to be resolvable by analyzing several mass range separately. Firstly, the loss rate of  
94 OVOC by condensation is strongly coupled to the molecular mass (Peräkylä et al., 2020), likely  
95 giving very different behaviors for the high and low mass ranges, even when produced by the same  
96 source. Second, dimers are a product of two RO<sub>2</sub>, which can have different sources, meaning that  
97 they may have temporal profiles unlike anything observable for monomers. Finally, if one mass range  
98 contains much less signal than another, it will have very little impact on the final PMF results.

99 In this study, we applied PMF analysis on three different mass ranges of mass spectra of OVOC  
100 measured by a chemical ionization atmospheric pressure interface time-of-flight (CI-APi-TOF,  
101 Jokinen et al. (2012)) mass spectrometer in the Finnish boreal forest. We utilized our recently

102 proposed new PMF approach, binPMF, to include as much of the high-resolution information in the  
103 mass spectra as possible, in a robust way (Zhang et al., 2019). We show the benefits of the sub-range  
104 PMF approach to better separate chemical sources, by reducing disturbance from variable loss terms  
105 of the OVOC. Much of the analysis focuses on dimer formation pathways, and the role of different  
106 nitrogen oxides in these pathways. We find that both daytime dimers and dimers resulting from the  
107 combination of different oxidants can be separated with the sub-range approach, but not with the PMF  
108 applied to the full mass range. We believe that this study will provide new perspectives for future  
109 studies analyzing gas-phase CIMS data.

110

## 111 **2 Methodology**

112 The focus of this work is on retrieving new information from mass spectra by applying new analytical  
113 approaches. Therefore, we chose a dataset that has been presented earlier, though without PMF  
114 analysis, by Zha et al. (2018), and was also used in the first study describing the binPMF method  
115 (Zhang et al., 2019). The measurements are described in more details below in section 2.1, while the  
116 data analysis techniques used in this work are presented in section 2.2.

117

### 118 **2.1 Measurements**

#### 119 **2.1.1 Ambient site**

120 The ambient measurements were conducted at the Station for Measuring Ecosystem–Atmosphere  
121 Relations (SMEAR) II in Finland (Hari and Kulmala, 2005) as part of the Influence of Biosphere–  
122 Atmosphere Interactions on the Reactive Nitrogen budget (IBAIRN) campaign (Zha et al, 2018).  
123 Located in the boreal environment in Hyytiälä, SMEAR II is surrounded with coniferous forest and  
124 has limited anthropogenic emission sources nearby. Diverse measurements of meteorology, aerosol  
125 and gas phase properties are continuously conducted at the station. Details about the meteorological  
126 conditions and temporal variations of trace gases during IBAIRN campaign are presented by Zha et  
127 al. (2018) and Liebmann et al. (2018).

128

#### 129 **2.1.2 Instrument and data**

130 Data were collected with a nitrate ( $\text{NO}_3^-$ )-based chemical ionization atmospheric pressure interface  
131 time-of-flight mass spectrometer (CI-APi-TOF, Jokinen et al., 2012) with about 4000 Th  $\text{Th}^{-1}$  mass  
132 resolving power, at ground level in September, 2016. In our study, the mass spectra were averaged to  
133 1 h time resolution from September 6<sup>th</sup> to 22<sup>nd</sup> for further analysis. We use the thomson (Th) as the  
134 unit for mass/charge, with 1 Th = 1 Da/e, where  $e$  is the elementary charge. As all the data discussed  
135 in this work are based on negative ion mass spectrometry, we will use the absolute value of the

136 mass/charge, although the charge of each ion will be negative. The masses discussed in this work  
137 includes the contribution from the nitrate ion, 62, unless specifically mentioned. Furthermore, as the  
138 technique is based on soft ionization with  $\text{NO}_3^-$  ions, any multiple charging effects are unlikely, and  
139 therefore the reported mass/charge values in thomson can be considered equivalent to the mass of the  
140 ion in Da.

141 The forest site of Hyytiälä is dominated by monoterpene emissions (Hakola et al., 2006). The main  
142 feature of previous CI-API-TOF measurements in Hyytiälä (Ehn et al., 2014; Yan et al., 2016) has  
143 been a bimodal distributions of HOM, termed monomers and dimers, as they are formed of either one  
144 or two  $\text{RO}_2$  radicals, respectively. For the analysis in this study, we chose three mass/charge ( $m/z$ )  
145 ranges of 50 Th each (Figure 1), corresponding to regions between which we expect differences in  
146 formation or loss mechanisms. In addition to regions with HOM monomers and HOM dimers, one  
147 range was chosen at lower masses, in a region presumably mainly consisting of molecules that are  
148 less likely to condense onto aerosol particles (Peräkylä et al., 2020).

149

## 150 **2.2 Positive matrix factorization (PMF)**

151 After the model of PMF was developed (Paatero and Tapper, 1994), numerous applications have been  
152 conducted with different types of environmental data (Song et al., 2007; Ulbrich et al., 2009; Yan et  
153 al., 2016; Zhang et al., 2017). By reducing dimensionality of the measured dataset, PMF model greatly  
154 simplifies the data analysis process with no requirement for prior knowledge of sources or pathways  
155 as essential input. The main factors can be further interpreted with their unique/dominant markers  
156 (elements or masses).

157 The basic assumption for PMF modelling is mass balance, which assumes that ambient concentration  
158 of a chemical component is the sum of contributions from several sources or processes, as shown in  
159 equation (1).

$$160 \quad \mathbf{X} = \mathbf{TS} \times \mathbf{MS} + \mathbf{R} \quad (1)$$

161 In equation (1),  $\mathbf{X}$  stands for the time series of measured concentration of different variables ( $m/z$  in  
162 our case),  $\mathbf{TS}$  represents the temporal variation of factor contributions,  $\mathbf{MS}$  stands for factor profiles  
163 (mass spectral profiles), and  $\mathbf{R}$  is the residual as the difference of the modelled and the observed data.  
164 The matrices  $\mathbf{TS}$  and  $\mathbf{MS}$  are iteratively calculated by a least-squares algorithm utilizing uncertainty  
165 estimates, to pursue minimized  $Q$  value as shown in equation (2), where  $S_{ij}$  is the estimated  
166 uncertainty, an essential input in PMF model.

$$167 \quad Q = \sum \sum \left( \frac{R_{ij}}{S_{ij}} \right)^2 \quad (2)$$

168 PMF model was conducted by multi-linear engine (ME-2) (Paatero, 1999) interfaced with Source  
169 Finder (SoFi, v6.3) (Canonaco et al., 2013). Signal-to-noise ratio (SNR) was calculated as  $SNR_{ij} =$   
170  $abs(X_{ij}) / abs(S_{ij})$ . When the Signal-to-noise ratio (SNR) is below 1, the signal of  $X_{ij}$  will be down-  
171 weighted by replacing the corresponding uncertainty  $S_{ij}$  by  $S_{ij}/SNR_{ij}$  (Visser et al., 2015). Future  
172 studies should pay attention to the potential risk when utilizing this method since down-weighting  
173 low signals element-wise will create a positive bias to the data. Robust mode was operated in the  
174 PMF modelling, where outliers ( $\left|\frac{R_{ij}}{S_{ij}}\right| > 4$ ) were significantly down-weighted (Paatero, 1997).

175

## 176 **2.3 binPMF**

177 As a newly developed application of PMF for mass spectral data, binPMF has no requirement for  
178 chemical composition information, while still taking advantage of the HR mass spectra, saving effort  
179 and time (Zhang et al., 2019). To explore the benefits of analyzing separated mass ranges, we applied  
180 binPMF to the three separated ranges. The three ranges were also later combined for binPMF analysis  
181 as comparison with the previous results. The PMF model requires both data matrix and error matrix  
182 as input, and details of the preparation of data and error matrices are described below.

183

### 184 **2.3.1 Data matrix**

185 Different from normal UMR or HR peak fitting, in binPMF, the mass spectra are divided into small  
186 bins after baseline subtraction and mass axis calibration. Linear interpolation was first conducted to  
187 the mass spectra with a mass interval of 0.001 Th. Then the interpolated data was averaged into bins  
188 of 0.02 Th width. We selected three ranges for further analysis based on earlier studies (Ehn et al.,  
189 2014; Yan et al., 2016; Bianchi et al., 2019; Peräkylä et al., 2020).

- 190 - Range 1,  $m/z$  250 – 300 Th, 51 unit masses  $\times$  25 bins per unit mass = 1275 bins/variables,  
191 consisting mainly of molecules with five to nine carbon atoms and four to nine oxygen atoms  
192 in our dataset.
- 193 - Range 2,  $m/z$  300 – 350 Th,  $51 \times 25 = 1275$  bins, mainly corresponding to HOM monomer  
194 products, featured with nine to ten C- and seven to ten O-atoms.
- 195 - Range 3,  $m/z$  510 – 560 Th,  $51 \times 30 = 1530$  bins, mainly corresponding to HOM dimer products,  
196 with carbon numbers of sixteen to twenty and eleven to fifteen O-atoms.

197 To avoid unnecessary computation, only signal regions with meaningful signals in the mass spectra  
198 were binned (Zhang et al., 2019). For a nominal mass  $N$ , the signal region included in further analyses  
199 was between  $N-0.2$  Th and  $N+0.3$  Th for Range 1 and 2, and between  $N-0.2$  Th and  $N+0.4$  Th for

200 Range 3. The wider signal regions in Range 3 is due to wider peaks at higher masses. The data were  
201 averaged into 1-h time resolution and in total we had 384 time points in the data matrix.

202

### 203 **2.3.2 Error matrix**

204 The error matrix represents the estimated uncertainty for each element of the data matrix and is crucial  
205 for iterative calculation of the  $Q$  minimum. Equation (3) is used for error estimation (Polissar et al.,  
206 1998),

$$207 \quad S_{ij} = \sigma_{ij} + \sigma_{\text{noise}} \quad (3)$$

208 where  $S_{ij}$  represents the uncertainty of  $m/z, j$  at time  $i$ ,  $\sigma_{ij}$  stands for counting statistics uncertainty  
209 and is estimated as follows,

$$210 \quad \sigma_{ij} = a \times \frac{\sqrt{I_{ij}}}{\sqrt{t_s}} \quad (4)$$

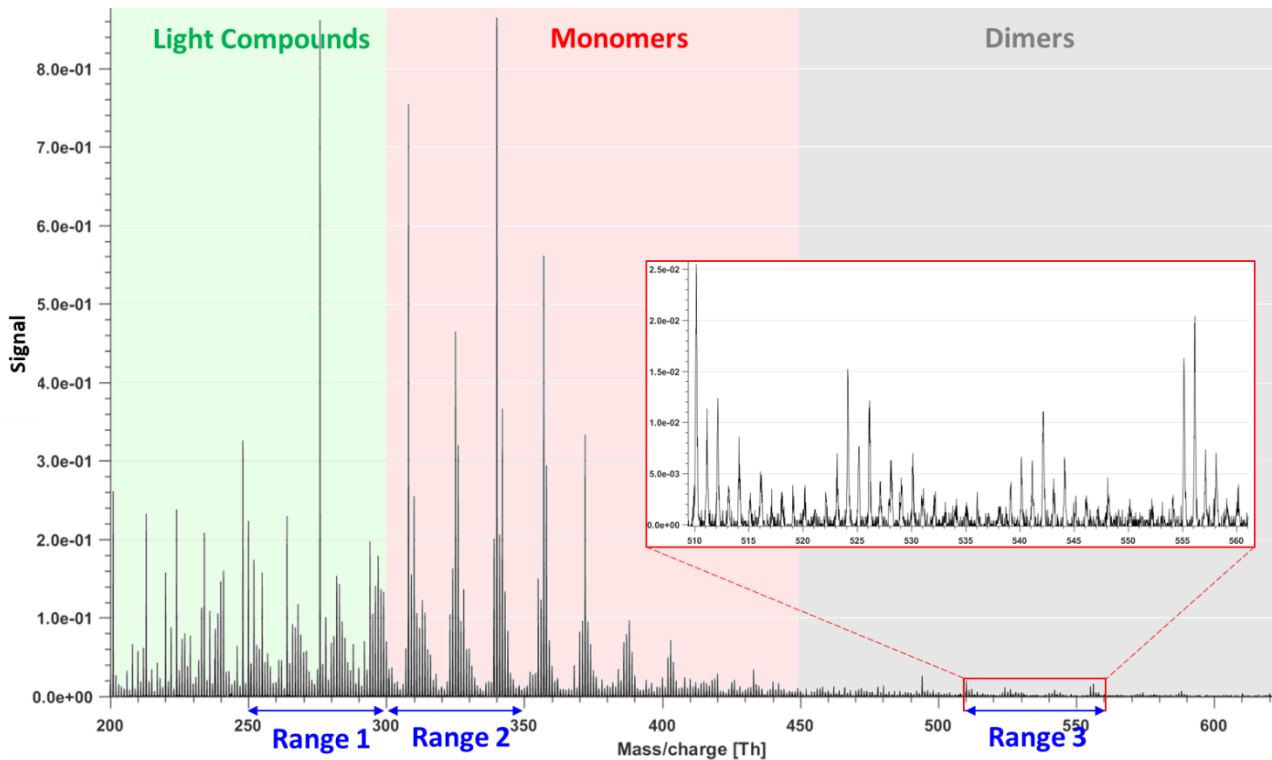
211 where  $I$  is the signal intensity term, in unit of counts per second (cps),  $t_s$  stands for length of averaging  
212 in seconds, while  $a$  is an empirical coefficient to compensate for unaccounted uncertainties (Allan et  
213 al., 2003; Yan et al., 2016) and is 1.28 in our study as previously estimated from laboratory  
214 experiments (Yan et al., 2016). The  $\sigma_{\text{noise}}$  term was estimated as the median of the standard  
215 deviations from signals in the bins in the region between nominal masses, where no physically  
216 meaningful signals are expected.

217

## 218 **3 Results**

### 219 **3.1 General overview of the dataset/spectrum**

220 During the campaign, in autumn, 2016, the weather was overall sunny and humid with average  
221 temperature of 10.8 °C and relative humidity (RH) of 87% (Zha et al., 2019). The average  
222 concentration of  $\text{NO}_x$  and  $\text{O}_3$  was 0.4 ppbv and 21 ppbv, respectively. The average total HOM  
223 concentration was  $\sim 10^8$  molecules  $\text{cm}^{-3}$ .



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Figure 1. Example of mass spectrum with 1-h time resolution measured from a boreal forest environment during the IBairn campaign (at 18:00, Finnish local time, UTC+2). The mass spectrum was divided into three parts and three sub-ranges were chosen from different parts for further analysis in our study. The nitrate ion (62 Th) is included in the mass.

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Figure 1 shows the 1 h averaged mass spectrum taken at 18:00 on September 12, as an example of the analyzed dataset. In addition to exploring the benefits of this type of sub-range analysis in relation to different formation or loss pathways, separating into sub-ranges may also aid factor identification for low-signal regions. As shown in Figure 1, there is a difference of 1-2 orders of magnitude in the signal intensity between Range 3 and Ranges 1-2. If all Ranges are run together, we would expect that the higher signals from Ranges 1 and 2 will drive the factorization. While if run separately, separating formation pathways of dimers in Range 3 will likely be easier. As dimers have been shown to be crucial for the formation of new aerosol particles from monoterpene oxidation (Kirkby et al., 2016; Troestl et al., 2016; Lehtipalo et al., 2018), this information may even be the most critical in some cases, despite the low contribution of these peaks to the total measured signal.

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binPMF was separately applied to Range 1, 2, 3, and a ‘Range combined’ which comprised all the three sub-ranges. All the PMF runs for the four ranges were conducted from two to ten factors and repeated three times for each factor number, to assure the consistency of the results. Factorization results and evolution with increasing factor number are briefly described in the following sections, separately for each Range (sections 3.2 – 3.5). It is worth noting that the factor order in factor



245 evolution does not necessarily correspond to that of the final results. The factor orders displayed in  
246 Figures 2-5 have been modified for further comparison between different ranges. More detailed  
247 discussion and comparison between the results are presented in Section 4.

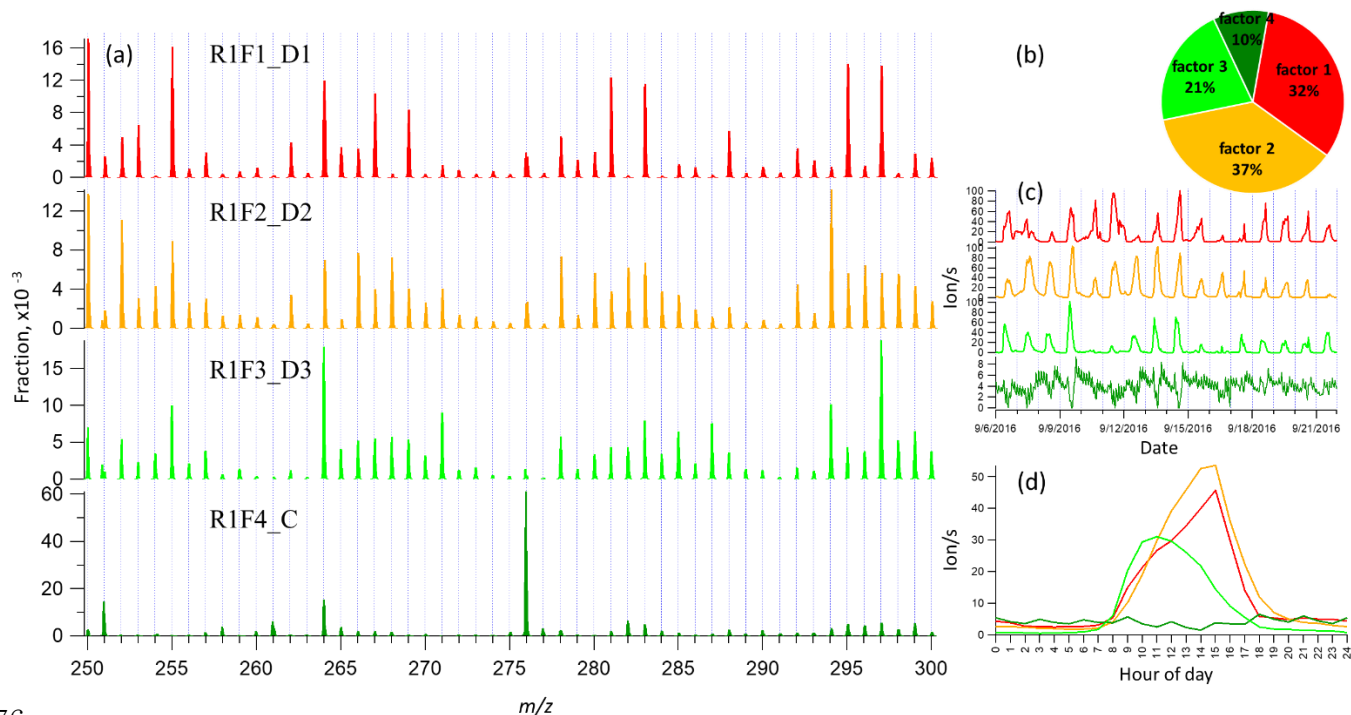
248

### 249 **3.2 binPMF on Range 1 (250 – 300 Th)**

250 As has become routine (Zhang et al., 2011; Craven et al., 2012), we first examined the mathematical  
251 parameters of our solutions. From two to ten factors,  $Q/Q_{exp}$  decreased from 2.8 to 0.7 (Fig S1 in  
252 supplementary information), and after three factors, the decreasing trend was gradually slowing down  
253 and approaching one, which is the ideal value for  $Q/Q_{exp}$  as a diagnostic parameter. The unexplained  
254 variation showed a decline from 18% to 8% from two to ten factors.

255 In the two-factor results, two daytime factors were separated, with peak time both at 14:00 - 15:00.  
256 One factor was characterized by large signals at 250 Th, 255 Th, 264 Th, 281 Th, 283 Th, 295 Th,  
257 297 Th. The other factor was characterized by large signals at 294 Th, 250 Th, 252 Th, 264 Th, 266  
258 Th, 268 Th, and 297 Th. In Hyytiälä, as reported in previous studies, odd masses observed by the  
259 nitrate CI-APi-TOF are generally linked to monoterpene-derived organonitrates during the day (Ehn  
260 et al., 2014; Yan et al., 2016). When the factor number increased to three, the two earlier daytime  
261 factors remained similar with the previous result, while a new factor appeared with a distinct sawtooth  
262 shape in the diurnal cycle. The main marker in the spectral profile was 276 Th, with a clear negative  
263 mass defect. When one more factor was added, the previous three factors remained similar as in the  
264 three-factor solution, and a new morning factor was resolved, with 264 Th and 297 Th dominant in  
265 the mass spectral profile, and a diurnal peak at 11:00.

266 As the factor number was increased, more daytime factors were separated, with similar spectral  
267 profiles to existing daytime factors and various peak times. No nighttime factors were found in the  
268 analysis even when the factor number reached ten. We chose the four-factor result for further  
269 discussion, and Figure 2 shows the result of Range 1, with spectral profile, time series, diurnal cycle  
270 and averaged factor contribution during the campaign. As shown in Figure 2d, Factors 1-3 are all  
271 daytime factors, while Factor 4 has no clear diurnal cycle, but a distinct sawtooth shape. Factor 4  
272 comes from a contamination of perfluorinated acids, from the inlet's automated zeroing every three  
273 hours during the measurements (Zhang et al., 2019). The zeroing periods have been removed from  
274 the dataset before binPMF analysis, but the contamination factor was still resolved. This factor is  
275 discussed in more detail in sections 4.1 and 4.4.



276

277 Figure 2 Four-factor result for Range 1, for (a) factor spectral profiles, (b) averaged factor contribution during  
 278 the campaign, (c) time series and (d) diurnal trend. Details on the factors' naming schemes are shown in Table  
 279 1.

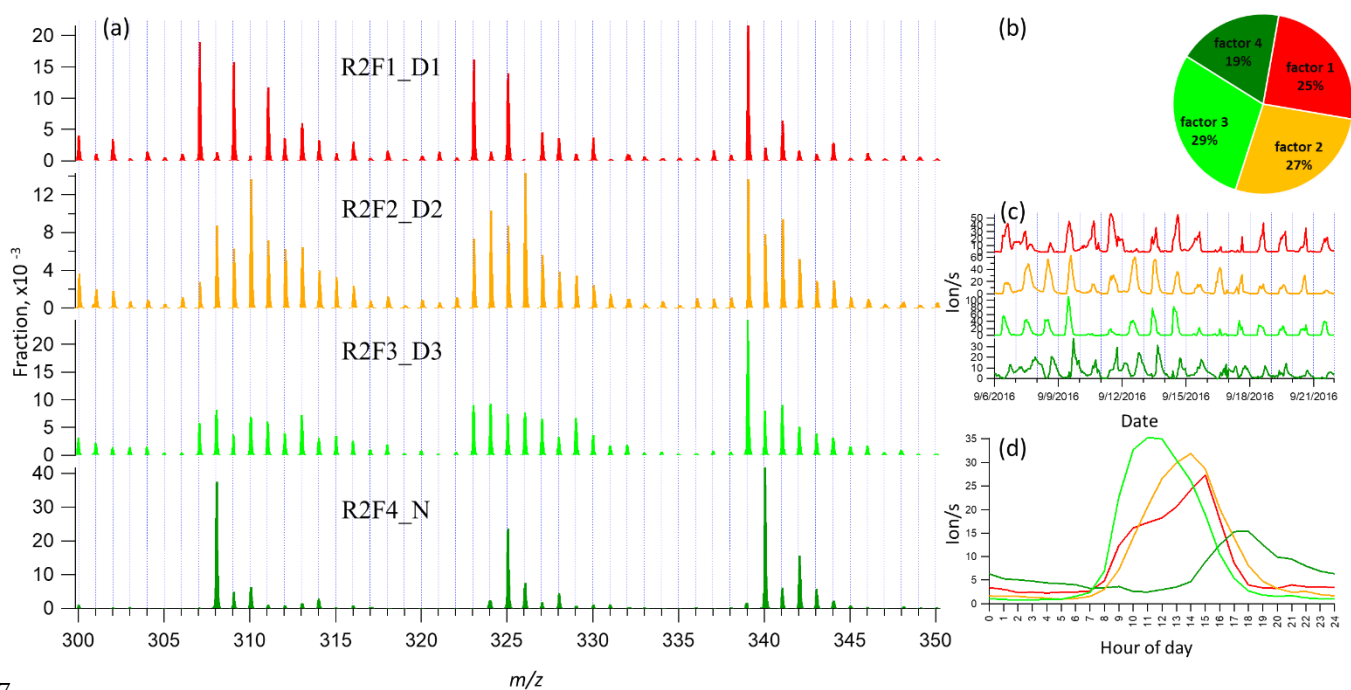
280

### 281 3.3 binPMF on Range 2 (300-350 Th)

282 This range covers the monoterpene HOM monomer range, and binPMF results have already been  
 283 discussed by Zhang et al. (2019) as a first example of the application of binPMF on ambient data.  
 284 Our input data here is slightly different. In the previous study, the 10 min automatic zeroing every  
 285 three hours was not removed before averaging to 1 hour time resolution while here, we have removed  
 286 this data. Overall, the results are similar as in our earlier study, and therefore the result are just briefly  
 287 summarized below for further comparison and discussion in Section 4. Similar to Range 1, both the  
 288  $Q/Q_{exp}$  (2.2 to 0.6) and unexplained variation (16% to 8%) declined with the increased factor number  
 289 from two to ten.

290 When the factor number was two, one daytime and one nighttime factor were separated, with diurnal  
 291 peak times at 14:00 and 17:00, respectively. The nighttime factor was characterized by masses at 340  
 292 Th, 308 Th and 325 Th (monoterpene ozonolysis HOM monomers (Ehn et al., 2014)) and remained  
 293 stable throughout the factor evolution from two to ten factors. With the addition of more factors, no  
 294 more nighttime factors got separated while the daytime factor was further separated and more daytime  
 295 factors appeared, peaking at various times in the morning (10:00 am), at noon or in the early afternoon  
 296 (around 14:00 pm and 15:00 pm). High contribution of 339 Th can be found in all the daytime factor

297 profiles. As the factor number reached six, a contamination factor appeared, characterized by large  
 298 signals at 339 Th and 324 Th, showing negative mass defects (Figure S2 in the Supplement). The  
 299 factor profile is nearly identical to the contamination factor determined in Zhang et al. (2019), where  
 300 the zeroing periods were not removed, causing larger signals for the contaminants. In our dataset,  
 301 where the zeroing periods were removed, no sawtooth pattern was discernible in the diurnal trend,  
 302 yet it could still be separated even though it only contributed 3% to Range 2. More about the  
 303 contamination factors from different ranges will be discussed in Section 4.4. We chose to show the  
 304 four-factor result below, to simplify the later discussion and comparison. Figure 3 shows four-factor  
 305 result of Range 2, with spectral profile, time series, diurnal cycle and averaged factor contribution  
 306 during the campaign.

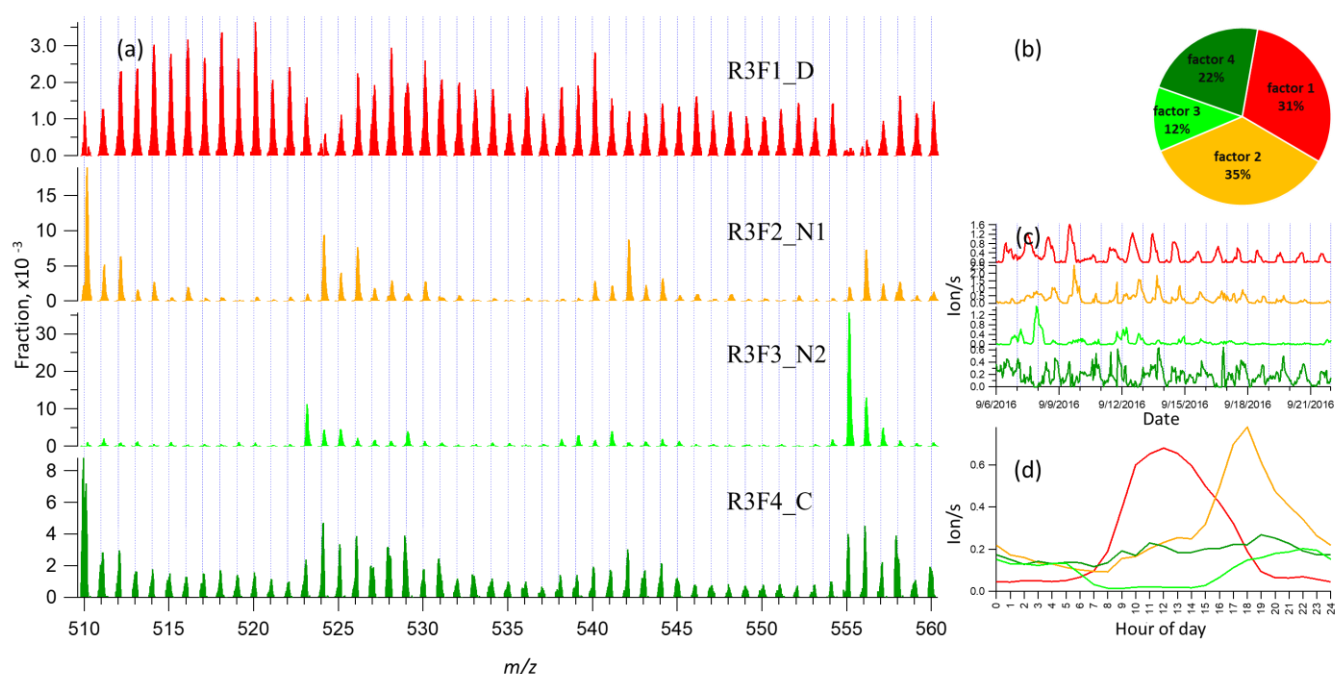


307  
 308 Figure 3 Four-factor result for Range 2, for (a) factor spectral profiles, (b) averaged factor contribution during  
 309 the campaign, (c) time series and (d) diurnal trend. Details on the factors' naming schemes are shown in Table  
 310 1.

311  
 312 **3.4 binPMF on Range 3 (510-560 Th)**

313 Range 3 represents mainly the monoterpene HOM dimers (Ehn et al., 2014). Similar to Range 1 and  
 314 2, both the  $Q/Q_{exp}$  (1.5 to 0.6) and unexplained variation (18% to 15%) showed decreasing trend with  
 315 the increased factor number (2-10). As can be seen from Figure 1, data in Range 3 had much lower  
 316 signals, compared to that of the Range 1 and 2, explaining the higher unexplained variation for Range  
 317 3.

318 In the two-factor result for Range 3, one daytime and one nighttime factor appeared, with diurnal  
 319 peak times at noon and 18:00, respectively. The nighttime factor was characterized by ions at 510 Th,  
 320 524 Th, 526 Th, 542 Th, and 555 Th, 556 Th, while the daytime factor showed no dominant marker  
 321 masses, yet with relatively high signals at 516 Th, 518 Th and 520 Th. As the number of factors  
 322 increased to three, one factor with almost flat diurnal trend was separated, with dominant masses of  
 323 510 Th, 529 Th, 558 Th. Most peaks in this factor had negative mass defects, and this factor was  
 324 again linked to a contamination factor. The four-factor result resolved another nighttime factor with  
 325 a dominant peak at 555 Th, and effectively zero contribution during daytime. As the factor number  
 326 was further increased, the new factors seemed like splits from previous factors with similar spectral  
 327 profiles. We therefore chose four-factor result also for Range 3 (results shown in Fig. 4) for further  
 328 discussion.

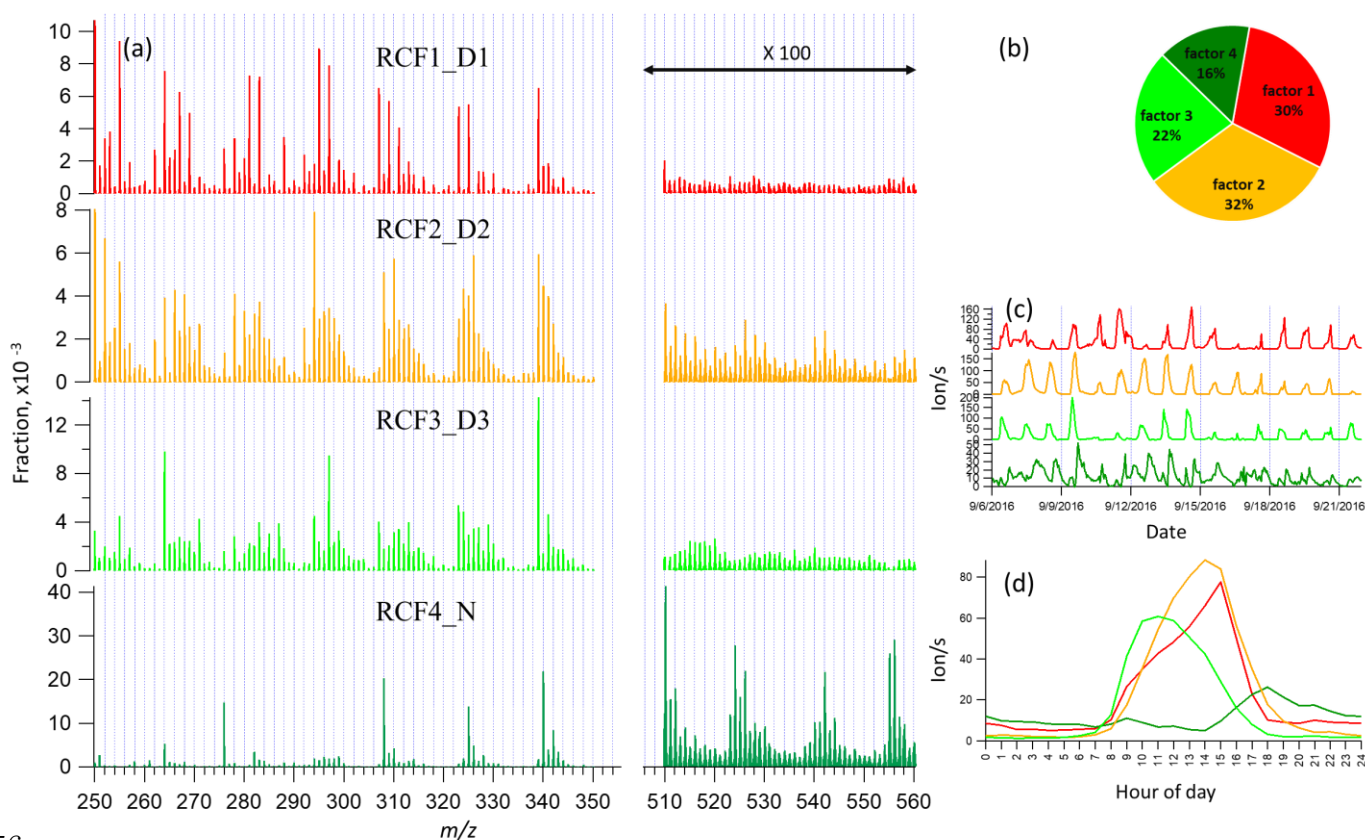


329  
 330 Figure 4 Four-factor result for Range 3, for (a) factor spectral profiles, (b) averaged factor contribution during  
 331 the campaign, (c) time series and (d) diurnal trend. Details on the factors' naming schemes are shown in Table  
 332 1.

333  
 334 **3.5 binPMF on Range Combined (250-350 Th & 510-560 Th)**

335 As comparison to the previous three ranges, we conducted the binPMF analysis on Range Combined,  
 336 which is the combination of the three ranges. The results of this range are fairly similar to those of  
 337 Ranges 1 and 2, as could be expected since the signal intensities in these ranges were much higher  
 338 than in Range 3. As the number of factors increased (2-10), both the  $Q/Q_{exp}$  (1.3 to 0.6) and  
 339 unexplained variation (16% to 8%) showed a decreasing trend.

340 In the two-factor result, one daytime factor and one nighttime factor were separated. In the nighttime  
 341 factor, most masses were found at even masses, and the fraction of masses in Range 3 was much  
 342 higher than that in daytime factor. In contrast, in the daytime factor, most masses were observed at  
 343 odd masses and the fraction of signal in Range 3 was much lower. During the day, photochemical  
 344 reactions as well as potential emissions increase the concentration of NO, which serves as peroxy  
 345 radical (RO<sub>2</sub>) terminator and often outcompetes RO<sub>2</sub> cross reactions in which dimers can be formed  
 346 (Ehn et al., 2014). Thus, the production of dimers is suppressed during the day, yielding instead a  
 347 larger fraction of organic nitrates, as has been shown also previously (Yan et al., 2016).  
 348 With the increase of the number of factors, more daytime factors were resolved with different peak  
 349 times. When the factor number reached seven, a clear sawtooth-shape diurnal cycle occurred, i.e. the  
 350 contamination factor, caused by the zeroing. As more factors were added, no further nighttime factors  
 351 were separated, and only more daytime factors appeared. To simplify the discussion and inter-range  
 352 comparison, we also here chose the four-factor result for further analysis. Figure 5 shows the four-  
 353 factor result of Range Combined, with spectral profile, time series, diurnal cycle and averaged factor  
 354 contribution during the campaign. The signals in range of 510-560 Th were enlarged 100-fold to be  
 355 visible.



356

357 Figure 5 Four-factor result for Range Combined, for (a) factor spectral profiles, (b) averaged factor  
 358 contribution during the campaign, (c) time series and (d) diurnal trend. Details on the factors' naming  
 359 schemes are shown in Table 1.

360

#### 361 4 Discussion

362 In Section 3, results by binPMF analysis were shown for Ranges 1, 2, 3 and Combined. In this section,  
 363 we discuss and compare the results from the different ranges. To simplify the inter-range comparison,  
 364 we chose four-factor results for all the four ranges, with the abbreviations shown in Table 1. From  
 365 Range 1, three daytime factors and a contaminations factor were separated. In Range 2, three daytime  
 366 factors and one nighttime factor (abbreviated as R2F4\_N) were resolved. The R2F4\_N factor was  
 367 characterized by signals at 308 Th ( $C_{10}H_{14}O_7 \cdot NO_3^-$ ), 325 Th ( $C_{10}H_{15}O_8 \cdot NO_3^-$ ), and 340 Th  
 368 ( $C_{10}H_{14}O_9 \cdot NO_3^-$ ), and can be confirmed as monoterpene ozonolysis products (Ehn et al., 2014; Yan et  
 369 al., 2016). With the increase of factor number to six, the contamination factor got separated also in  
 370 this mass range. In Range 3, one daytime factor, two nighttime factors and a contamination factor  
 371 were separated. The first nighttime factor (R3F2\_N1) had large peaks at 510 Th ( $C_{20}H_{32}O_{11} \cdot NO_3^-$ )  
 372 and 556 Th ( $C_{20}H_{30}O_{14} \cdot NO_3^-$ ), dimer products that have been identified during chamber studies of  
 373 monoterpene ozonolysis (Ehn et al., 2014). The molecule observed at 510 Th has 32 H-atoms,  
 374 suggesting that one of the  $RO_2$  involved would have been initiated by OH, which is formed during  
 375 the ozonolysis of alkenes such as monoterpenes at nighttime (Atkinson et al., 1992; Paulson and  
 376 Orlando, 1996). The other nighttime factor (R3F3\_N2) was dominated by ions at 523 Th  
 377 ( $C_{20}H_{31}O_8NO_3 \cdot NO_3^-$ ) and 555 Th ( $C_{20}H_{31}O_{10}NO_3 \cdot NO_3^-$ ), representing nighttime monoterpene  
 378 oxidation involving  $NO_3$ . As these dimers contain only one N-atom, and 31 H-atoms, we can assume  
 379 that they are formed from reactions between an  $RO_2$  formed from  $NO_3$  oxidation and another  $RO_2$   
 380 formed by ozone oxidation. These results match well with the profiles in a previous study by Yan et  
 381 al. (2016). The results of Range Combined are very similar to Range 2, with one nighttime factor and  
 382 three daytime factors. The contamination factor was separated with increase of factor number to seven.

383

384

Table 1. Summary of PMF results for the different mass ranges

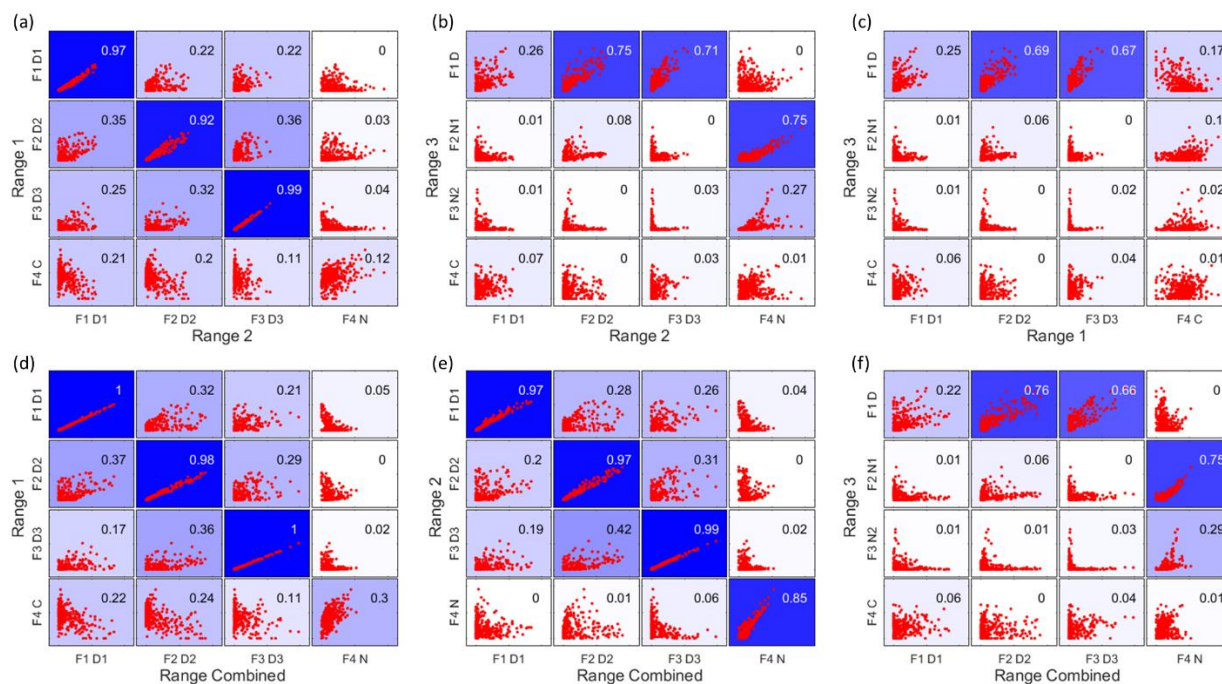
Range	Factor number	Factor name <sup>a</sup>	Dominant peaks	Peak time
1 (250 - 300 Th)	1	R1F1_D1	250, 255, 295, 297	15:00
	2	R1F2_D2	250, 252, 294	15:00
	3	R1F3_D3	264, 297	11:00
	4	R1F4_C	276	<sup>b</sup>
2 (300 - 350 Th)	1	R2F1_D1	307, 309, 323, 325, 339,	15:00
	2	R2F2_D2	310, 326, 339,	14:00
	3	R2F3_D3	339	11:00
	4	R2F4_N	308, 325, 340	18:00

<b>3 (510 – 560 Th)</b>	1	R3F1_D	516, 518, 520, 528, 540	12:00
	2	R3F2_N1	510, 524, 542, 556	18:00
	3	R3F3_N2	523, 555	22:00
	4	R3F4_C	510, 558	- <sup>b</sup>
<b>Combined (1, 2, 3)</b>	1	RCF1_D1	250, 255, 295, 339	15:00
	2	RCF2_D2	250, 252, 294, 339	14:00
	3	RCF3_D3	264, 297, 339	11:00
	4	RCF4_N	308, 340, 510, 524, 555, 556	18:00

385 <sup>a</sup> Factor name is defined with range name, factor number and name. For example, RxFy represents Factor y in Range x.  
386 RC stands for Range Combined. For the factor name, D is short for daytime, N for Nighttime, C for contamination.  
387 <sup>b</sup> The contamination factor in Range 1 shows sawtooth pattern; while in Range 3 shows no diurnal pattern.  
388

#### 389 4.1 Time series correlation

390 In Figure 6, the upper panels show the time series correlations among the first three ranges. As  
391 expected based on the results above, generally the daytime factors, and the two nighttime  
392 monoterpene ozonolysis factors (R2F4\_N and R3F2\_N1) correlated well, respectively. However, the  
393 contamination factors did not show strong correlation between different ranges, even though  
394 undoubtedly from the same source. More about the contamination factors will be discussed in Section  
395 4.4. The lower panels in Figure 6 displays the correlations between the first three ranges and the  
396 Range Combined, and clearly demonstrates that the results of Range Combined is mainly controlled  
397 by high signals from Range 1 and 2. More detailed aspects of the comparison between factors in  
398 different ranges is given in the following sections. The good agreements between factors from  
399 different subranges also help to verify the robustness of the solutions.



400  
401 Figure 6 Time series correlations among Range 1, 2, 3 (upper panels a-c), and between the first three ranges  
402 and the Range Combined (lower panels d-f). The abbreviations for different factors are the same in Table 1,  
403 with F for factor, D for daytime, N for nighttime and C for contamination, e.g. F1D1 for Factor 1 daytime 1.

404 The coefficient of determination,  $R^2$ , is marked in each subplot by a number shown in the right upper corners  
405 and by the blue colors, with stronger blue indicating higher  $R^2$ .

406

## 407 **4.2 Daytime processes**

### 408 **4.2.1 Factor comparison**

409 As mentioned above, with increasing number of factors, usually more daytime factors will be resolved,  
410 reflecting the complicated daytime photochemistry. The three daytime factors between Range 1 and  
411 2 agreed with each other quite well (Figure 6a). However, R1F1\_D1 and R2F1\_D1 did not show  
412 strong correlation with the only daytime factor in Range 3 (R3F1\_D), while the other two daytime  
413 factors in both Range 1 and 2, i.e. R1F2\_D2, R1F3\_D3, R2F2\_D2, R2F3\_D3, correlated well with  
414 R3F1\_D from Range 3.

415 The 1<sup>st</sup> daytime factors from Range 1 and 2, R1F1\_D1 and R2F1\_D1, were mainly characterized by  
416 odd masses 255 Th, 281 Th, 283 Th, 295 Th, 297 Th, 307 Th, 309 Th, 311 Th, 323 Th, 325 Th, 339  
417 Th. The factors are dominated by organonitrates. Organic nitrate formation during daytime is  
418 generally associated with the termination of RO<sub>2</sub> radicals by NO. This termination step is mutually  
419 exclusive with the termination of RO<sub>2</sub> with other RO<sub>2</sub>, which can lead to dimer formation. If the NO  
420 concentration is the limiting factor for the formation of these factors, the low correlations between  
421 the NO-terminated monomer factors, and the dimer factors, is to be expected. In contrast, if the other  
422 daytime factors mainly depend on oxidant and monoterpene concentrations, some correlation  
423 between those, and the daytime dimer factor, is to be expected, as shown in Figure 6b, c.

424 All the spectral profiles resolved from Range Combined binPMF analysis inevitably contained mass  
425 contribution from 510 – 560 Th, even the daytime factor from Range Combined (RCF1\_D1) which  
426 did not show clear correlation with R3F1\_D from Range 3 (Figure 6e).

427 The 2<sup>nd</sup> and 3<sup>rd</sup> daytime factors in Range 1 and 2, R1F2\_D2, R1F3\_D3, R2F2\_D2, R2F3\_D3, had  
428 high correlations with R3F1\_D in Range 3. Daytime factors in Range Combined (RCF2\_D2 and  
429 RCF3\_D3) also showed good correlation with R3F1\_D in Range 3. However, if we compare R3F1\_D  
430 and the mass range of 510 – 560 Th of the daytime factors in Range Combined, just with a quick look,  
431 we can readily see the difference. The daytime factor separated in Range 3 (R3F1\_D) has no obvious  
432 markers in the profile. With the increase of factor number (up to ten factors), no clearly new factors  
433 were separated in Range 3, but instead the previously separated factors were seen to split into several  
434 factors. However, the spectral pattern in R3F1\_D is different from that in the mass range of 510 –  
435 560 Th in RCF2\_D2. The factorization of Range Combined was mainly controlled by low masses  
436 due to their high signals. The signals at high masses were forced to be distributed according to the



437 time series determined by small masses. Ultimately, this will lead to failure in factor separation for  
 438 this low-signal range.

#### 439 4.2.2 Daytime dimer formation

440 Dimers are primarily produced during nighttime, due to NO suppressing  $RO_2 + RO_2$  reactions in  
 441 daytime (Ehn et al., 2014; Yan et al., 2016). However, in this study, we found one clear daytime factor  
 442 in Range 3 (R3F1\_D, peak at local time 12:00, UTC+2) by sub-range analysis. With high loadings  
 443 from even masses including 516, 518, 520, 528, 540 Th, this only daytime factor in dimer range  
 444 correlated very well with two daytime factors in Ranges 1 and 2 (R1F2\_D2, R1F3\_D3, R2F2\_D2,  
 445 R2F3\_D3) (Figure 6b and c). Table 2 include the correlation matrix of all PMF and factors and  
 446 selected meteorological parameters. Strong correlation between R3F1\_D with solar radiation was  
 447 found, with  $R = 0.79$  (Table 2). This may indicate involvement of OH oxidation in producing this  
 448 factor.

449 Table 2 Correlation between factors and meteorological parameters and gases

	R1F1_D1	R1F1_D2	R1F1_D3	R1F1_C	R2F1_D1	R2F2_D2	R2F3_D3	R2F4_N	R3F1_D	R3F2_N1	R3F3_N2	R3F4_C	RCF1_D1	RCF2_D2	RCF3_D3	RCF4_N
O <sub>3</sub>	0.51	0.59	0.35	-0.18	0.47	0.57	0.36	0.43	0.55	0.33	0.27	0.22	0.49	0.57	0.33	0.34
NO	0.13	-0.01	0.24	-0.03	0.18	-0.02	0.24	-0.22	0.13	-0.19	-0.17	0.03	0.13	0.00	0.26	-0.18
NO <sub>x</sub>	-0.05	-0.22	-0.10	0.09	-0.01	-0.23	-0.11	-0.13	-0.16	-0.21	-0.04	0.04	-0.04	-0.22	-0.09	-0.11
RH	-0.46	-0.80	-0.63	0.30	-0.43	-0.82	-0.64	-0.27	-0.78	-0.39	-0.07	-0.07	-0.43	-0.82	-0.60	-0.21
T	0.66	0.72	0.40	-0.24	0.65	0.66	0.41	0.39	0.65	0.30	0.14	0.19	0.66	0.68	0.38	0.24
UVB	0.52	0.63	0.82	-0.40	0.52	0.68	0.84	-0.30	0.79	-0.08	-0.27	0.08	0.49	0.68	0.83	-0.29

450  
 451 As previous studies have shown, dimers greatly facilitate new particle formation (NPF) (Kirkby et  
 452 al., 2016; Troestl et al., 2016; Lehtipalo et al., 2018), and this daytime dimer factor may represent a  
 453 source of dimers that would impact the initial stages of NPF in Hyytiälä. Mohr et al. (2017) reported  
 454 a clear diel pattern of dimers (sum of about 60 dimeric compounds of  $C_{16-20}H_{13-33}O_{6-9}$ ) during NPF  
 455 events in 2013 in Hyytiälä, with minimum at night and maximum after noon, and estimated these  
 456 dimers can contribute ~5% of the mass of sub-60 nm particles. The link between the dimers presented  
 457 in that paper and those reported here will require further studies, as will the proper quantification of  
 458 the dimer factor identified here.

#### 459 4.3 Nighttime processes

##### 460 4.3.1 Factor comparison

461 Since high-mass dimers are more likely to form at night due to photochemical production of NO in  
 462 daytime, which inhibits  $RO_2 + RO_2$  reactions, Range 3 had the highest fraction of nighttime signals  
 463

464 of all the sub-ranges. While Range 3 produced two nighttime factors, Ranges 2 and Combined showed  
465 one, and Range 1 had no nighttime factor. The difference between the two results also indicates the  
466 advantage of analyzing monomers and dimers separately.

467 The two nighttime factors in Range 3 can be clearly identified as arising from ozonolysis (R3F2\_N1)  
468 and a mix of ozonolysis and NO<sub>3</sub> oxidation (R3F2\_N2) based on the mass spectral profiles, as  
469 described above. The organonitrate at 555 Th, C<sub>20</sub>H<sub>31</sub>O<sub>10</sub>NO<sub>3</sub>·NO<sub>3</sub><sup>-</sup>, is a typical marker for NO<sub>3</sub>  
470 radical initiated monoterpene chemistry (Yan et al., 2016). However, several interesting features  
471 become evident when comparing to the results of Range 2 and Combined. Firstly, only one nighttime  
472 factor (R2F4\_N, RCF4\_N) was separated in each of these ranges, and that shows clear resemblance  
473 with ozonolysis of monoterpenes as measured in numerous studies, e.g. Ehn et al. (2012);(2014).  
474 Secondly, the high correlation found in Figure 6b between the ozonolysis factors (i.e., R2F4\_N,  
475 R3F2\_N1, RCF4\_N), further supports the assignment. However, factor R2F4\_N is the only nighttime  
476 factor in the monomer range, suggesting that NO<sub>3</sub> radical chemistry of monoterpenes in Hyytiälä does  
477 not form substantial amounts of HOM monomers. The only way for the CI-APi-TOF to detect  
478 products of monoterpene-NO<sub>3</sub> radical chemistry may thus be through the dimers, where one highly  
479 oxygenated RO<sub>2</sub> radical from ozonolysis reacts with a less oxygenated RO<sub>2</sub> radical from NO<sub>3</sub>  
480 oxidation.

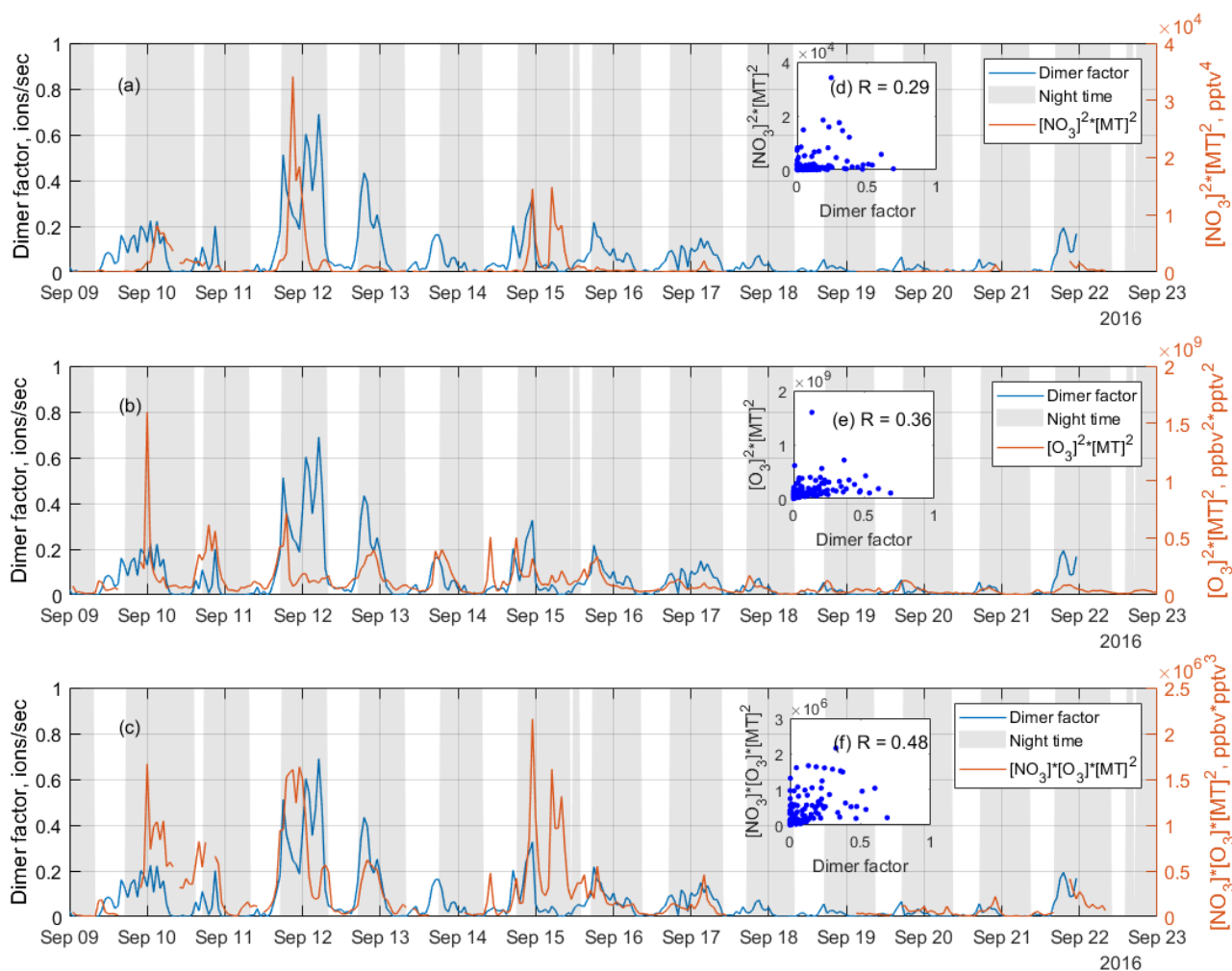
481 In the results by Yan et al. (2016) the combined UMR-PMF of monomers and dimers did yield a  
482 considerable amount of compounds in the monomer range also for the NO<sub>3</sub> radical chemistry factor.  
483 There may be several reasons for this discrepancy. One major cause for differences between the spring  
484 dataset of Yan et al. (2016) and the autumn dataset presented here, is that nighttime concentrations  
485 of HOM was greatly reduced during our autumn campaign. The cause may have been fairly frequent  
486 fog formation during nights, and also the concentration of e.g. ozone decreased nearly to zero during  
487 several nights (Zha et al., 2018). It is also possible that the NO<sub>3</sub> radical-related factor by Yan et al.  
488 (2016) is probably a mixture of NO<sub>3</sub> and O<sub>3</sub> radical chemistry, while the monomer may thus be  
489 attributed to the O<sub>3</sub> part. Alternatively, the different conditions during the two measurement periods,  
490 as well as seasonal difference in monoterpene mixtures (Hakola et al., 2012), caused variations in the  
491 oxidation pathways.

492

#### 493 **4.3.2 Dimers initiated by NO<sub>3</sub> radicals**

494 Previous studies show that NO<sub>3</sub> oxidation of α-pinene, the most abundant monoterpene in Hyytiälä  
495 (Hakola et al., 2012), produces fairly little SOA mass (yields <4 %), while β-pinene shows yields of  
496 up to 53 % (Bonn and Moorgat, 2002;Nah et al., 2016). The NO<sub>3</sub>+β-pinene reaction results in low  
497 volatile organic nitrate compounds with carboxylic acid, alcohol, and peroxide functional groups (Fry

498 et al., 2014;Boyd et al., 2015), while  $\text{NO}_3 + \alpha\text{-pinene}$  reaction will typically lose the nitrate functional  
 499 group and form oxidation products with high vapor pressures (Spittler et al., 2006;Perraud et al.,  
 500 2010). Most monoterpene-derived HOM, including monomers, are low-volatile (Peräkylä et al.,  
 501 2020) and thus a low SOA yield indicates a low HOM yield. Thus, while there are to our knowledge  
 502 no laboratory studies on HOM formation from  $\text{NO}_3$  oxidation of  $\alpha\text{-pinene}$ , a low yield can be expected  
 503 based on SOA studies.



504

505 Figure 7 Time series of the  $\text{NO}_3$  oxidation dimer factor (blue line), and the product of (a)  $[\text{NO}_3]^2 \times$   
 506  $[\text{monoterpene}]^2$ , (b)  $[\text{O}_3]^2 \times [\text{monoterpene}]^2$ , and (c)  $[\text{NO}_3] \times [\text{O}_3] \times [\text{monoterpene}]^2$ , where []  
 507 represents concentration in unit of pptv for  $\text{NO}_3$  radicals and monoterpene, ppbv for  $\text{O}_3$ , while the scatter plots are shown  
 508 as inserts, (d), (e), (f), respectively. The scatter plots and correlation coefficients R are only calculated from  
 509 nighttime data, which is selected based on solar radiation, to eliminate the influence from daytime oxidation  
 510 processes.

511  
 512 As discussed above, a dimer factor (R3F2\_N2) was identified as being a crossover between  $\text{NO}_3$   
 513 radical initiated and  $\text{O}_3$  initiated  $\text{RO}_2$  radicals. Figure 7 shows the time series of this factor, as well as  
 514 the product of  $[\text{NO}_3]^2 \times [\text{monoterpene}]^2$ ,  $[\text{O}_3]^2 \times [\text{monoterpene}]^2$ , and  $[\text{NO}_3] \times [\text{O}_3] \times [\text{monoterpene}]^2$ .  
 515 These products are used to mimic the formation rates of the  $\text{RO}_2$  radicals reacting to form the dimers,

516 either from pure NO<sub>3</sub> oxidation (Fig. 7a), pure O<sub>3</sub> oxidation (7b), or the mixed reaction between RO<sub>2</sub>  
517 from the two oxidants (7c). The NO<sub>3</sub> concentration was estimated in Liebmann et al. (2018) for the  
518 same campaign. Monoterpenes were measured using a proton transfer reaction time of flight mass  
519 spectrometer (PTR-TOF-MS). More details on measurement of NO<sub>3</sub> proxy and monoterpene can be  
520 found in in Liebmann et al. (2018).

521 As shown in Figure 7, the time series of the dimer factor tracks those of [NO<sub>3</sub>] × [monoterpene] and  
522 [O<sub>3</sub>] × [monoterpene] reasonably well, but shows the highest correlation with the product of [NO<sub>3</sub>] ×  
523 [O<sub>3</sub>] × [monoterpene]<sup>2</sup>. This further supports this dimer formation as a mixed processes of ozonolysis  
524 and NO<sub>3</sub> oxidation. The heterogeneity of the monoterpene emissions in the forest, and the fact that  
525 no dimer loss process is included, partly explain the relatively low correlation coefficients. The  
526 sampling inlets for PTR-TOF were about 170 m away from the NO<sub>3</sub> reactivity measurement  
527 (Liebmann et al., 2018), which in turn was some tens of meters away from the HOM measurements.  
528 Thus, this analysis should be considered qualitative only.

529 The nitrate dimer factor (R3F2\_N2) was dominated by the organonitrate at 555 Th,  
530 C<sub>20</sub>H<sub>31</sub>O<sub>10</sub>NO<sub>3</sub>·NO<sub>3</sub><sup>-</sup>. However, unlike the pure ozonolysis dimer factor which had a corresponding  
531 monomer factor (R = 0.86 between factor R2F4\_N and R3 F2\_N1), this NO<sub>3</sub>-related dimer factor did  
532 not have an equivalent monomer factor. This suggests that the NO<sub>3</sub> oxidation of the monoterpene  
533 mixture in Hyytiälä does not by itself form much HOM, but in the presence of RO<sub>2</sub> from ozonolysis,  
534 the RO<sub>2</sub> from NO<sub>3</sub> oxidation can take part in HOM dimer formation. This further implies that,  
535 different from previous knowledge based on single-oxidant experiments in chambers, NO<sub>3</sub> oxidation  
536 may have a larger impact on SOA formation in the atmosphere where different oxidants exist  
537 concurrently. This highlights the need for future laboratory studies to consider systems with multiple  
538 oxidants during monoterpene oxidation experiments, to truly understand the role and contribution of  
539 different oxidants, and NO<sub>3</sub> in particular.

#### 540 **4.4 Fluorinated compounds**

541 During the campaign, an automated instrument zeroing every three hours was conducted. While the  
542 zeroing successfully removed the low-volatile HOM and H<sub>2</sub>SO<sub>4</sub>, the process also introduced  
543 contaminants into the inlet lines, e.g. perfluorinated organic acids from Teflon tubing. Each zeroing  
544 process lasted for 10 min. In the data analysis, we removed all the 10-min zeroing periods, and  
545 averaged the data to 1-h time resolution, but contaminants were still identified in all ranges by  
546 binPMF. However, the correlation between contamination factors from different ranges is low (Figure  
547 6c).

548 To future investigate the low factor correlations of the same source, three fluorinated compounds with  
549 different volatilities,  $(\text{CF}_2)_3\text{CO}_2\text{HF}\cdot\text{NO}_3^-$  (275.9748 Th),  $(\text{CF}_2)_5\text{C}_2\text{O}_4\text{H}$  (338.9721 Th), and  
550  $(\text{CF}_2)_6\text{CO}_2\text{HF}\cdot\text{NO}_3^-$  (425.9653 Th), were examined in fine time resolution, i.e. 1 min. The time series  
551 and 3-h cycle of the three fluorinated compounds were shown in Figure S3 and S4 in Supplement.  
552 The correlation coefficients dropped greatly before and after the zero period was removed, from 0.9  
553 to 0.3 for  $R^2$  between 276 Th and 339 Th, and 0.8 to 0.1 between 276 Th and 426 Th (Fig. S5a, b).  
554 Similar effect is also found with the 1 h averaged data (Fig. S5c, d). It is evident that the three  
555 fluorinated compounds were from the same source (zeroing process), but due to their different  
556 volatilities, they were lost at different rates. This, in turn, means that the spectral signature of this  
557 source will change as a function of time, at odds with one of the basic assumptions of PMF.  
558 The analysis of the fluorinated compounds in our system was here merely used as an example to show  
559 that volatility can impact source profiles over time. In Figure S5, it can be clearly seen that the profile  
560 of Range Combined is noisier than that of Range 3, probably due to the varied fractional contributions  
561 of contamination compounds to the profile. In ambient data, products from different sources can have  
562 undergone atmospheric processing, altering the product distribution. This analysis highlighted the  
563 importance of differences in the sink terms due to different volatilities of the products. This may be  
564 an important issue for gas phase mass spectrometry analysis, potentially underestimated by many  
565 PMF users, as it is likely only a minor issue for aerosol data, for which PMF has been applied much  
566 more routinely. If failing to achieve physically meaningful factors using PMF on gas phase mass  
567 spectra, our recommendation is to try applying PMF to sub-ranges of the spectrum, where IVOC,  
568 SVOC and (E)LVOC could be analyzed separately.

569

#### 570 **4.5 Atmospheric insights**

571 Based on the new data analysis technique binPMF applied in sub-ranges of mass spectra, we were  
572 able to separate two particularly intriguing atmospheric processes, the formation of daytime dimers  
573 as well as dimer formation involving  $\text{NO}_3$  radicals, which otherwise could not have been identified  
574 in our study.

575 With a diurnal peak around noon time, the daytime dimers identified in this study correlate very well  
576 with daytime factors in monomer range. Strong correlation between this factor and solar radiation  
577 indicate the potential role of OH oxidation in the formation of daytime dimers. By now, very few  
578 studies have reported the observations of daytime dimers. As dimers are shown to be able to take part  
579 in new particle formation (NPF) (Kirkby et al., 2016), this daytime dimer may contribute to the early  
580 stages of NPF in the boreal forest.

581 The second process identified in our study is the formation of dimers that are a crossover between  
582  $\text{NO}_3$  and  $\text{O}_3$  oxidation. Such dimers have been identified before (Yan et al., 2016). However, we were  
583 not able to identify corresponding HOM monomer compounds. This finding indicates that while  $\text{NO}_3$   
584 oxidation of the monoterpenes in Hyytiälä may not undergo autoxidation to form HOM by themselves,  
585 they can contribute to HOM dimers when the  $\text{NO}_3$ -derived  $\text{RO}_2$  react with highly oxygenated  $\text{RO}_2$   
586 from other oxidants. Multi-oxidant systems should be taken into consideration in future experimental  
587 studies on monoterpene oxidation processes.

588

## 589 **5 Conclusions**

590 The recent developments in the field of mass spectrometry, combined with factor analysis techniques  
591 such as PMF, have greatly improved our understanding of complicated atmospheric processes and  
592 sources. In this study, we applied the new binPMF approach (Zhang et al., 2019), to separate sub-  
593 ranges of mass spectra measured using a chemical ionization mass spectrometer in the Finnish boreal  
594 forest. By using this method, we were able to identify a daytime dimer factor, presumably initiated  
595 by  $\text{OH}/\text{O}_3$  oxidation of monoterpenes, forming from  $\text{RO}_2+\text{RO}_2$  reactions despite competition from  
596 daytime  $\text{NO}$ . This compound group, showing a diurnal peak around noon, may contribute to new  
597 particle formation at the site. In addition, we successfully separated  $\text{NO}_3$ -related dimers which would  
598 not have been identified from this dataset without utilizing the different sub-ranges. The  $\text{NO}_3$ -related  
599 factor was consistent with earlier observations (Yan et al., 2016), with the exception that we did not  
600 observe any corresponding monomer factor. This may be explained by the observed nitrate-  
601 containing dimers being formed from two  $\text{RO}_2$ , where one is initiated by oxidation by  $\text{O}_3$ , and the  
602 other by  $\text{NO}_3$ . If the  $\text{NO}_3$ -derived  $\text{RO}_2$  are not able to form HOM by themselves, there will not be any  
603 related monomers observed. To validate this hypothesis, future laboratory experiments that target  
604 more complex oxidation systems will be useful in order to understand the role of  $\text{NO}_3$  oxidation in  
605 SOA formation under different atmospheric conditions.

606 Apart from these two major findings, we also found several other benefits of applying PMF on  
607 separate sub-ranges of the mass spectra. First, different compounds from the same source can have  
608 variable loss rates due to differences in volatilities. This leads to increased difficulty for PMF to  
609 separate this source, but if the PMF analysis is run separately on lighter masses (with higher volatility)  
610 and heavier masses (with lower volatility), the source may become easier to distinguish. Secondly,  
611 chemistry or sources contributing only to one particular mass range, e.g. dimers, can be better  
612 separated. Thirdly, mass ranges with small, but informative, signals can be more accurately assigned  
613 as their contribution becomes larger than if the entire mass range was analyzed at once. Finally,  
614 running PMF on separate mass ranges also allows comparing the factors between the different ranges,

615 helping to verify the results. In summary, while we do not suggest that this type of sub-range analysis  
616 should always be utilized, we recommend other analysts of gas-phase mass spectrometer data to test  
617 this approach in order to see whether additional useful information can be obtained. In our dataset,  
618 this method was crucial for identifying different types of dimers and dimer formation pathways,  
619 which are of great importance for the formation of both new particles and SOA.

620

626 **Data availability.** The data used in this study are available from the first author upon request: please  
627 contact Yanjun Zhang (yanjun.zhang@helsinki.fi).

635 **Author contribution.** ME and YZ designed the study. QZ and MR collected the data; data analysis  
636 and manuscript writing were done by YZ. All coauthors discussed the results and commented the  
637 manuscript.

638 **Competing interests.** The authors declare that they have no conflict of interest

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