

1 **Insights on Atmospheric Oxidation Processes by Performing Factor Analyses on** 2 **Sub-ranges of Mass Spectra**

3 Yanjun Zhang¹, Otso Peräkylä¹, Chao Yan¹, Liine Heikkinen¹, Mikko Äijälä¹, Kaspar R. Daellenbach¹,
4 Qiaozhi Zha¹, Matthieu Riva^{1,2}, Olga Garmash¹, Heikki Junninen^{1,3}, Pentti Paatero¹, Douglas Worsnop^{1,4}, and
5 Mikael Ehn¹

6 ¹ Institute for Atmospheric and Earth System Research / Physics, Faculty of Science, University of Helsinki,
7 Helsinki, 00014, Finland

8 ² Univ Lyon, Université Claude Bernard Lyon 1, CNRS, IRCELYON, F-69626, Villeurbanne, France

9 ³ Institute of Physics, University of Tartu, Tartu, 50090, Estonia

10 ⁴ Aerodyne Research, Inc., Billerica, MA 01821, USA

11
12 Corresponding author: yanjun.zhang@helsinki.fi

13 14 **Abstract**

15 With the recent developments in mass spectrometry, combined with the strengths of factor analysis
16 techniques, our understanding of atmospheric oxidation chemistry has improved significantly. The
17 typical approach for using techniques like positive matrix factorization (PMF) is to input all measured
18 data for the factorization in order to separate contributions from different sources and/or processes to
19 the total measured signal. However, while this is a valid approach for assigning the total signal to
20 factors, we have identified several cases where useful information can be lost if solely using this
21 approach. For example, gaseous molecules emitted from the same source can show different temporal
22 behaviors due to differing loss terms, like condensation at different rates due to different molecular
23 masses. This conflicts with one of PMF's basic assumptions of constant factor profiles. In addition,
24 some ranges of a mass spectrum may contain useful information, despite contributing only minimal
25 fraction to the total signal, in which case they are unlikely to have a significant impact on the
26 factorization result. Finally, certain mass ranges may contain molecules formed via pathways not
27 available to molecules in other mass ranges, e.g. dimeric species versus monomeric species. In this
28 study, we attempted to address these challenges by dividing mass spectra into sub-ranges and
29 applying the newly developed binPMF method to these ranges separately. We utilized a dataset from
30 a chemical ionization atmospheric pressure interface time-of-flight (CI-APi-TOF) mass spectrometer
31 as an example. We compare the results from these three different ranges, each corresponding to
32 molecules of different volatilities, with binPMF results from the combined range. Separate analysis
33 showed clear benefits in dividing factors for molecules of different volatilities more accurately, in

34 resolving different chemical processes from different ranges, and in giving a chance for high-
35 molecular-weight molecules with low signal intensities to be used to distinguish dimeric species with
36 different formation pathways. As two major insights from our study, we identified daytime dimer
37 formation (diurnal peak around noon) which may contribute to new particle formation in Hyytiälä, as
38 well as dimers from NO₃ oxidation process. We recommend PMF users to try running their analyses
39 on selected sub-ranges in order to further explore their datasets.

41 **1 Introduction**

42 Huge amounts of volatile organic compounds (VOC) are emitted to the atmosphere every year
43 (Guenther et al., 1995;Lamarque et al., 2010), which play a significant role in atmospheric chemistry
44 and affect the oxidative ability of the atmosphere. The oxidation products of VOC can contribute to
45 the formation and growth of secondary organic aerosols (Kulmala et al., 2013;Ehn et al., 2014;Kirkby
46 et al., 2016;Troestl et al., 2016), affecting air quality, human health, and climate radiative forcing
47 (Pope III et al., 2009;Stocker et al., 2013;Zhang et al., 2016;Shiraiwa et al., 2017). Thanks to the
48 advancement in mass spectrometric applications, like the aerosol mass spectrometer (AMS)
49 (Canagaratna et al., 2007) and chemical ionization mass spectrometry (CIMS) (Bertram et al.,
50 2011;Jokinen et al., 2012;Lee et al., 2014) our capability to detect these oxidized products, as well as
51 our understanding of the complicated atmospheric oxidation pathways in which they take part, have
52 been greatly enhanced.

53 Monoterpenes (C₁₀H₁₆), one common group of VOC emitted in forested areas, have been shown to
54 be a large source of atmospheric secondary organic aerosol (SOA). The oxidation of monoterpenes
55 produces a wealth of different oxidation products (Oxygenated VOC, OVOC), including highly
56 oxygenated organic molecules (HOM) with molar yields in the range of a few percent, depending on
57 the specific monoterpene and oxidant (Ehn et al., 2014;Bianchi et al., 2019). Bianchi et al. (2019)
58 summarized that HOM can be either Extremely Low Volatility Organic Compounds (ELVOC), Low
59 Volatility Organic Compounds (LVOC), or Semi-volatile Organic Compounds (SVOC)
60 (classifications by Donahue et al. 2012), depending on their exact structures. For less oxygenated
61 products, the majority are likely to fall into the SVOC or the Intermediate VOC (IVOC) range. The
62 volatility of the OVOC will determine their dynamics, including their ability to contribute to the
63 formation of SOA and new particles (Bianchi et al., 2019;Buchholz et al., 2019).

64 The recent developments of CIMS techniques has allowed researchers to observe unprecedented
65 numbers of OVOC, in real-time (Riva et al., 2019). This ability to measure thousands of compounds
66 is a great benefit, but also a large challenge for the data analyst. For this reason, factor analytical
67 techniques have often been applied to reduce the complexity of the data by finding co-varying signals

68 that can be grouped into common factors (Huang et al., 1999). For aerosol and gas-phase mass
69 spectrometry, positive matrix factorization, PMF (Paatero and Tapper, 1994;Zhang et al., 2011) has
70 been the most utilized tool. The factors have then been attributed to sources (e.g. biomass burning
71 organic aerosol) or processes (e.g. monoterpene ozonolysis) depending on the application and ability
72 to identify spectral signatures (Yan et al., 2016;Zhang et al., 2017). In the vast majority of these PMF
73 applications to mass spectra, the mass range of ions has been maximized in order to provide as much
74 input as possible for the algorithm. This approach was certainly motivated in early application of
75 PMF on e.g. offline filters, with chemical information of metals, water-soluble ions, and organic and
76 elemental carbon (OC and EC), where the number of variables is counted in tens, and the number of
77 samples in tens or hundreds (Zhang et al., 2017). However, with gas-phase CIMS, we often have up
78 to a thousand variables, with hundreds or even thousands of samples, meaning that the amount of data
79 itself is unlikely to be a limitation for PMF calculation. In this work, we aimed to explore potential
80 benefits of dividing the spectra into sub-ranges before applying factorization analysis.

81 An inherent requirement of factorization approaches is that the factor profiles, in this case the relative
82 abundancies of ions in the mass spectra, of each factor stay nearly constant. Due to the complexity
83 and number of atmospheric processes affecting the formation, transformation, and loss of VOC,
84 OVOC and aerosol, this often does not hold, and is one of the main limitations of factorization
85 approaches. Given the different volatilities of OVOC, it may even be expected that molecules from
86 the same source may have very different loss time scales, which may affect the factor analysis. For
87 compounds of low volatility, such as many HOM, the main atmospheric loss process is typically
88 condensation onto aerosol particles, with chemical sink being negligible (Bianchi et al., 2019). If, on
89 the other hand, a compound does not irreversibly condense, oxidation reactions can also affect its
90 lifetime. Volatility issue has been studied and reported for AMS data, with different volatilities of
91 various OA types (Huffman et al., 2009;Crippa et al., 2014;Paciga et al., 2016;Äijälä et al., 2017).
92 Semi-volatile oxygenated organic aerosol (SV-OOA) and Low-volatility oxygenated organic aerosol
93 (LV-OOA) can both be mainly produced from biogenic sources, but get separated based on different
94 volatilities by PMF (El Haddad et al., 2013). Sekimoto et al. (2018) found that the two profiles
95 resolved with VOC emitted from biomass burning had different estimated volatilities. As the
96 volatility of a molecule is linked to its molecular mass (Peräkylä et al., 2020), it may be beneficial to
97 apply PMF separately to mass ranges where one can expect the loss processes to be similar, thereby
98 resulting in more constant factor profiles. In this way, distinct sources are hopefully separated by
99 PMF, with minimized influence of differing volatilities from one source.

100 The number of PMF or other factorization studies utilizing CIMS data remains very limited.
101 “Traditional” PMF analyses have so far, to our knowledge, only been applied to nitrate-based

102 chemical ionization atmospheric pressure interface time-of-flight (CI-APi-TOF) data (Yan et al.,
103 2016;Massoli et al., 2018). One study has also utilized non-negative matrix factorization (NNMF) to
104 look at diurnal trends of Iodide ToF-CIMS data (Lee et al., 2018). The lack of more studies utilizing
105 PMF, or other factorization techniques, on CIMS data is likely partly due to the complexity of the
106 data, with multiple overlapping ions hampering HR peak fitting (Zhang et al., 2019). In addition,
107 variable factor profiles may hamper PMF's ability to correctly separate the factors. The two CI-APi-
108 TOF studies utilized the nearly the entire measured spectrum (from around 200 Th to around 600 Th),
109 either in unit mass resolution (UMR) or high resolution (HR) peak fitting data (Yan et al.,
110 2016;Massoli et al., 2018). Massoli et al. (2018) estimated the volatility of the molecules they detected,
111 finding that all the six extracted factors had notable contributions from IVOC, SVOC and (E)LVOC.
112 These compound groups will have clearly different loss mechanisms, and thereby loss rates,
113 suggesting that variation in factor profiles is inevitable, even if the source was identical for all
114 molecules in the factor. We hypothesize that this effect further hampers the correct factorization, and
115 further that this effect can be reduced by dividing the spectra into separate ranges, with each sub-
116 range containing molecules with roughly similar loss mechanisms and rates.

117 As an additional motivation to separate different ranges from the mass spectrum, it is not only the
118 loss mechanisms, but also the formation pathways that may differ. For example, atmospheric
119 oxidation chemistry of organics is, to a large extent, the chemistry of peroxy radicals (RO_2) (Orlando
120 and Tyndall, 2012). These RO_2 are initiated by VOC reacting with oxidants like ozone, or the hydroxyl
121 (OH) or nitrate (NO_3) radicals, while their termination occurs mainly by bimolecular reactions with
122 NO , HO_2 and/or other RO_2 . Some product molecules can be formed from any of the three termination
123 pathways, while for example ROOR "dimers" can only be formed from $\text{RO}_2 + \text{RO}_2$ reactions (Berndt
124 et al., 2018a;Berndt et al., 2018b). This also means that there can be several different pathways to
125 form dimers from the same precursors VOC, by combining RO_2 formed from the same or different
126 oxidants. As an example of the latter, an ROOR dimer can contain one moiety produced from ozone
127 oxidation and another moiety from NO_3 oxidation (Yan et al., 2016). Thus, their concentration is
128 dependent on both the precursor VOC concentration, and the concentrations of both oxidants. Such a
129 molecule will not have a direct equivalent in any of the monomer products: even though monomers
130 can form from $\text{RO}_2 + \text{R}'\text{O}_2$ reactions, the products from RO_2 are not dependent on the source of the
131 $\text{R}'\text{O}_2$. This may complicate the identification of certain dimer factors by PMF if the entire spectrum
132 is analyzed at once, and therefore separation of the monomer and dimer products before the PMF
133 analysis could be advantageous.

134 Recently, we proposed a new PMF approach, binPMF, to simplify the analysis of mass spectral data
135 (Zhang et al., 2019). This method divides the mass spectrum into narrow bins, typically some tens of

136 bins per integer mass, depending on the mass resolving power of the instrument, before performing
137 PMF analyses. In this way, binPMF does not require any time-consuming, and potentially subjective
138 high resolution peak fitting, and can thus be utilized for data exploration at a very early stage of data
139 analysis. Data preparation is nearly as simple as in the case of UMR analysis, yet it utilizes much
140 more spectral information as it does not sum up signal over all ions at each integer mass. In addition
141 to saving time and effort in data analysis, the results are less sensitive to mass calibration fluctuations.
142 Finally, the binning also greatly increases the number of input variables, which has the advantage that
143 factor analysis with smaller mass ranges becomes more feasible, as more meaningful variation is
144 supplied to the algorithm.

145 We designed this study to explore the benefits of separate analysis of different mass ranges from mass
146 spectra. We used a previously published ambient dataset measured by a CI-API-TOF, and conducted
147 binPMF analysis with three different mass ranges, roughly corresponding to different volatility ranges.
148 We compared the results from the sub-range analyses with each other and with results from binPMF
149 run on the combined ranges. We found that dimers generated during daytime and dimers initiated by
150 NO₃ oxidation can be separated from our dataset by utilizing the sub-ranges, but not with the full
151 range. We believe that this study will provide new perspectives for future studies analyzing gas-phase
152 CIMS data.

153

154 **2 Methodology**

155 The focus of this work is on retrieving new information from mass spectra by applying new analytical
156 approaches. Therefore, we chose a dataset that has been presented earlier, though without PMF
157 analysis, by Zha et al. (2018), and was also used in the first study describing the binPMF method
158 (Zhang et al., 2019). The measurements are described in more details below in section 2.1, while the
159 data analysis techniques used in this work are presented in section 2.2.

160

161 **2.1 Measurements**

162 **2.1.1 Ambient site**

163 The ambient measurements were conducted at the Station for Measuring Ecosystem–Atmosphere
164 Relations (SMEAR) II in Finland (Hari and Kulmala, 2005) as part of the Influence of Biosphere–
165 Atmosphere Interactions on the Reactive Nitrogen budget (IBAIRN) campaign (Zha et al, 2018).
166 Located in the boreal environment in Hyytiälä, SMEAR II is surrounded with coniferous forest and
167 has limited anthropogenic emission sources nearby. Diverse measurements of meteorology, aerosol
168 and gas phase properties are continuously conducted at the station. Details about the meteorological

169 conditions and temporal variations of trace gases during IBairn campaign are presented by Zha et
170 al. (2018) and Liebmann et al. (2018).

171

172 **2.1.2 Instrument and data**

173 Data were collected with a nitrate (NO_3^-)-based chemical ionization atmospheric pressure interface
174 time-of-flight mass spectrometer (CI-APi-TOF, Jokinen et al., 2012) with about 4000 Th Th^{-1} mass
175 resolving power, at ground level in September, 2016. In our study, the mass spectra were averaged to
176 1 h time resolution from September 6th to 22nd for further analysis. We use the thomson (Th) as the
177 unit for mass/charge, with 1 Th = 1 Da/e, where e is the elementary charge. As all the data discussed
178 in this work are based on negative ion mass spectrometry, we will use the absolute value of the
179 mass/charge, although the charge of each ion will be negative. The masses discussed in this work
180 includes the contribution from the nitrate ion, 62, unless specifically mentioned. Furthermore, as the
181 technique is based on soft ionization with NO_3^- ions, any multiple charging effects are unlikely, and
182 therefore the reported mass/charge values in thomson can be considered equivalent to the mass of the
183 ion in Da.

184 The forest site of Hyytiälä is dominated by monoterpene emissions (Hakola et al., 2006). The main
185 feature of previous CI-APi-TOF measurements in Hyytiälä (Ehn et al., 2014; Yan et al., 2016) has
186 been a bimodal distributions of HOM, termed monomers and dimers, as they are formed of either one
187 or two RO_2 radicals, respectively. For the analysis in this study, we chose three mass/charge (m/z)
188 ranges of 50 Th each (Figure 1), corresponding to regions between which we expect differences in
189 formation or loss mechanisms. In addition to regions with HOM monomers and HOM dimers, one
190 range was chosen at lower masses, in a region presumably mainly consisting of molecules that are
191 less likely to condense onto aerosol particles (Peräkylä et al., 2020).

192

193 **2.2 Positive matrix factorization (PMF)**

194 After the model of PMF was developed (Paatero and Tapper, 1994), numerous applications have been
195 conducted with different types of environmental data (Song et al., 2007; Ulbrich et al., 2009; Yan et
196 al., 2016; Zhang et al., 2017). By reducing dimensionality of the measured dataset, PMF model greatly
197 simplifies the data analysis process with no requirement for prior knowledge of sources or pathways
198 as essential input. The main factors can be further interpreted with their unique/dominant markers
199 (elements or masses).

200 The basic assumption for PMF modelling is mass balance, which assumes that ambient concentration
201 of a chemical component is the sum of contributions from several sources or processes, as shown in
202 equation (1).

203
$$\mathbf{X} = \mathbf{TS} \times \mathbf{MS} + \mathbf{R} \quad (1)$$

204 In equation (1), \mathbf{X} stands for the time series of measured concentration of different variables (m/z in
205 our case), \mathbf{TS} represents the temporal variation of factor contributions, \mathbf{MS} stands for factor profiles
206 (mass spectral profiles), and \mathbf{R} is the residual as the difference of the modelled and the observed data.
207 The matrices \mathbf{TS} and \mathbf{MS} are iteratively calculated by a least-squares algorithm utilizing uncertainty
208 estimates, to pursue minimized Q value as shown in equation (2), where S_{ij} is the estimated
209 uncertainty, an essential input in PMF model.

210
$$Q = \sum \sum \left(\frac{R_{ij}}{S_{ij}} \right)^2 \quad (2)$$

211 PMF model was conducted by multi-linear engine (ME-2) (Paatero, 1999) interfaced with Source
212 Finder (SoFi, v6.3) (Canonaco et al., 2013). Signal-to-noise ratio (SNR) was calculated as $\text{SNR}_{ij} =$
213 $\text{abs}(X_{ij}) / \text{abs}(S_{ij})$. When the Signal-to-noise ratio (SNR) is below 1, the signal of X_{ij} will be down-
214 weighted by replacing the corresponding uncertainty S_{ij} by S_{ij}/SNR_{ij} (Visser et al., 2015). Future
215 studies should pay attention to the potential risk when utilizing this method since down-weighting
216 low signals element-wise will create a positive bias to the data. Robust mode was operated in the
217 PMF modelling, where outliers ($\left| \frac{R_{ij}}{S_{ij}} \right| > 4$) were significantly down-weighted (Paatero, 1997).

218

219 **2.3 binPMF**

220 As a newly developed application of PMF for mass spectral data, binPMF has no requirement for
221 chemical composition information, while still taking advantage of the HR mass spectra, saving effort
222 and time (Zhang et al., 2019). To explore the benefits of analyzing separated mass ranges, we applied
223 binPMF to the three separated ranges. The three ranges were also later combined for binPMF analysis
224 as comparison with the previous results. The PMF model requires both data matrix and error matrix
225 as input, and details of the preparation of data and error matrices are described below.

226

227 **2.3.1 Data matrix**

228 Different from normal UMR or HR peak fitting, in binPMF, the mass spectra are divided into small
229 bins after baseline subtraction and mass axis calibration. Linear interpolation was first conducted to
230 the mass spectra with a mass interval of 0.001 Th. Then the interpolated data was averaged into bins
231 of 0.02 Th width. We selected three ranges for further analysis based on earlier studies (Ehn et al.,
232 2014; Yan et al., 2016; Bianchi et al., 2019; Peräkylä et al., 2020).

- 233 - Range 1, m/z 250 – 300 Th, 51 unit masses \times 25 bins per unit mass = 1275 bins/variables,
 234 consisting mainly of molecules with five to nine carbon atoms and four to nine oxygen atoms
 235 in our dataset.
- 236 - Range 2, m/z 300 – 350 Th, $51 \times 25 = 1275$ bins, mainly corresponding to HOM monomer
 237 products, featured with nine to ten C- and seven to ten O-atoms.
- 238 - Range 3, m/z 510 – 560 Th, $51 \times 30 = 1530$ bins, mainly corresponding to HOM dimer products,
 239 with carbon numbers of sixteen to twenty and eleven to fifteen O-atoms.

240 To avoid unnecessary computation, only signal regions with meaningful signals in the mass spectra
 241 were binned (Zhang et al., 2019). For a nominal mass N , the signal region included in further analyses
 242 was between $N-0.2$ Th and $N+0.3$ Th for Range 1 and 2, and between $N-0.2$ Th and $N+0.4$ Th for
 243 Range 3. The wider signal regions in Range 3 is due to wider peaks at higher masses. The data were
 244 averaged into 1-h time resolution and in total we had 384 time points in the data matrix.

245

246 2.3.2 Error matrix

247 The error matrix represents the estimated uncertainty for each element of the data matrix and is crucial
 248 for iterative calculation of the Q minimum. Equation (3) is used for error estimation (Polissar et al.,
 249 1998),

$$250 \quad S_{ij} = \sigma_{ij} + \sigma_{noise} \quad (3)$$

251 where S_{ij} represents the uncertainty of m/z j at time i , σ_{ij} stands for counting statistics uncertainty
 252 and is estimated as follows,

$$253 \quad \sigma_{ij} = a \times \frac{\sqrt{I_{ij}}}{\sqrt{t_s}} \quad (4)$$

254 where I is the signal intensity term, in unit of counts per second (cps), t_s stands for length of averaging
 255 in seconds, while a is an empirical coefficient to compensate for unaccounted uncertainties (Allan et
 256 al., 2003; Yan et al., 2016) and is 1.28 in our study as previously estimated from laboratory
 257 experiments (Yan et al., 2016). The σ_{noise} term was estimated as the median of the standard
 258 deviations from signals in the bins in the region between nominal masses, where no physically
 259 meaningful signals are expected.

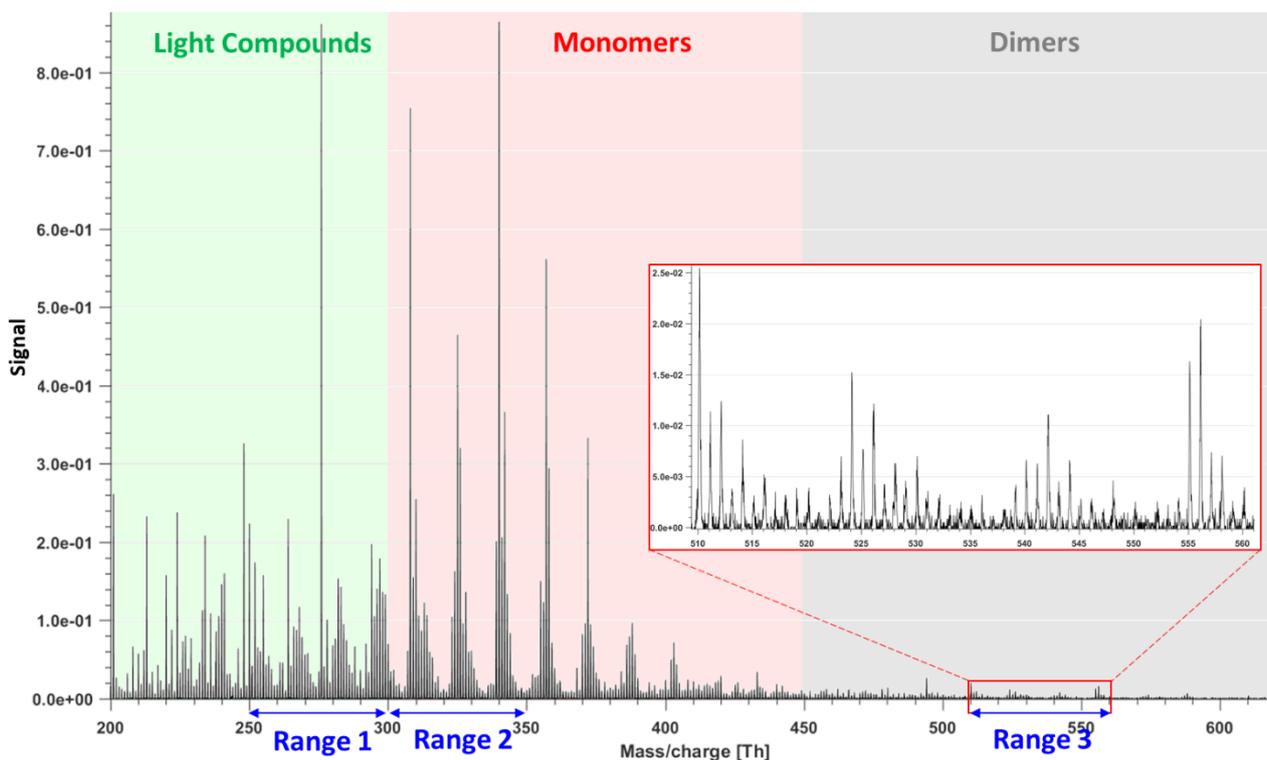
260

261 3 Results

262 3.1 General overview of the dataset/spectrum

263 During the campaign, in autumn, 2016, the weather was overall sunny and humid with average
 264 temperature of 10.8 °C and relative humidity (RH) of 87% (Zha et al., 2019). The average

265 concentration of NO_x and O₃ was 0.4 ppbv and 21 ppbv, respectively. The average total HOM
266 concentration was ~ 10⁸ molecules cm⁻³.



267
268 Figure 1. Example of mass spectrum with 1-h time resolution measured from a boreal forest environment
269 during the IBAIRN campaign (at 18:00, Finnish local time, UTC+2). The mass spectrum was divided into
270 three parts and three sub-ranges were chosen from different parts for further analysis in our study. The
271 nitrate ion (62 Th) is included in the mass.

272
273 Figure 1 shows the 1 h averaged mass spectrum taken at 18:00 on September 12, as an example of
274 the analyzed dataset. In addition to exploring the benefits of this type of sub-range analysis in relation
275 to different formation or loss pathways, separating into sub-ranges may also aid factor identification
276 for low-signal regions. As shown in Figure 1, there is a difference of 1-2 orders of magnitude in the
277 signal intensity between Range 3 and Ranges 1-2. If all Ranges are run together, we would expect
278 that the higher signals from Ranges 1 and 2 will drive the factorization. While if run separately,
279 separating formation pathways of dimers in Range 3 will likely be easier. As dimers have been shown
280 to be crucial for the formation of new aerosol particles from monoterpene oxidation (Kirkby et al.,
281 2016; Troestl et al., 2016; Lehtipalo et al., 2018), this information may even be the most critical in
282 some cases, despite the low contribution of these peaks to the total measured signal.

283 binPMF was separately applied to Range 1, 2, 3, and a 'Range combined' which comprised all the
284 three sub-ranges. All the PMF runs for the four ranges were conducted from two to ten factors and
285 repeated three times for each factor number, to assure the consistency of the results. Factorization

286 results and evolution with increasing factor number are briefly described in the following sections,
287 separately for each Range (sections 3.2 – 3.5). It is worth noting that the factor order in factor
288 evolution does not necessarily correspond to that of the final results. The factor orders displayed in
289 Figures 2-5 have been modified for further comparison between different ranges. More detailed
290 discussion and comparison between the results are presented in Section 4.

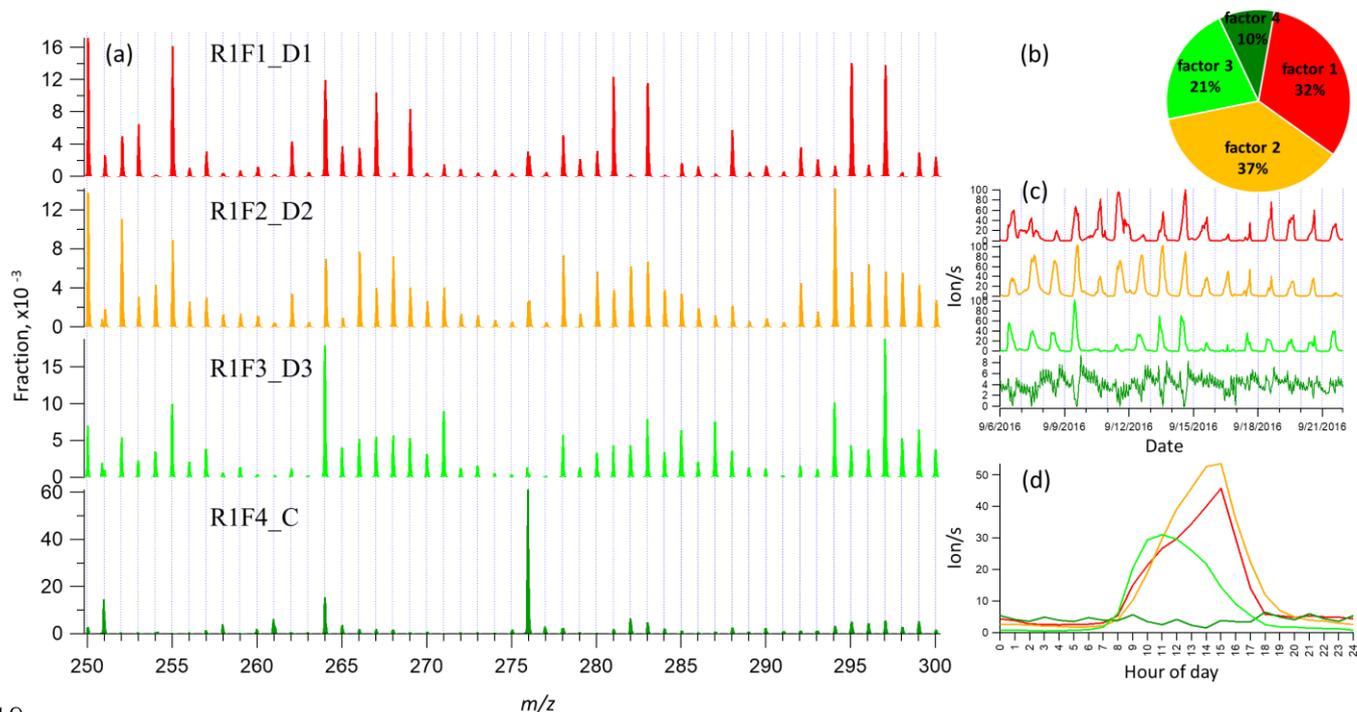
291

292 **3.2 binPMF on Range 1 (250 – 300 Th)**

293 As has become routine (Zhang et al., 2011; Craven et al., 2012), we first examined the mathematical
294 parameters of our solutions. From two to ten factors, Q/Q_{exp} decreased from 2.8 to 0.7 (Fig S1 in
295 supplementary information), and after three factors, the decreasing trend was gradually slowing down
296 and approaching one, which is the ideal value for Q/Q_{exp} as a diagnostic parameter. The unexplained
297 variation showed a decline from 18% to 8% from two to ten factors.

298 In the two-factor results, two daytime factors were separated, with peak time both at 14:00 - 15:00.
299 One factor was characterized by large signals at 250 Th, 255 Th, 264 Th, 281 Th, 283 Th, 295 Th,
300 297 Th. The other factor was characterized by large signals at 294 Th, 250 Th, 252 Th, 264 Th, 266
301 Th, 268 Th, and 297 Th. In Hyytiälä, as reported in previous studies, odd masses observed by the
302 nitrate CI-APi-TOF are generally linked to monoterpene-derived organonitrates during the day (Ehn
303 et al., 2014; Yan et al., 2016). When the factor number increased to three, the two earlier daytime
304 factors remained similar with the previous result, while a new factor appeared with a distinct sawtooth
305 shape in the diurnal cycle. The main marker in the spectral profile was 276 Th, with a clear negative
306 mass defect. When one more factor was added, the previous three factors remained similar as in the
307 three-factor solution, and a new morning factor was resolved, with 264 Th and 297 Th dominant in
308 the mass spectral profile, and a diurnal peak at 11:00.

309 As the factor number was increased, more daytime factors were separated, with similar spectral
310 profiles to existing daytime factors and various peak times. No nighttime factors were found in the
311 analysis even when the factor number reached ten. We chose the four-factor result for further
312 discussion, and Figure 2 shows the result of Range 1, with spectral profile, time series, diurnal cycle
313 and averaged factor contribution during the campaign. As shown in Figure 2d, Factors 1-3 are all
314 daytime factors, while Factor 4 has no clear diurnal cycle, but a distinct sawtooth shape. Factor 4
315 comes from a contamination of perfluorinated acids, from the inlet's automated zeroing every three
316 hours during the measurements (Zhang et al., 2019). The zeroing periods have been removed from
317 the dataset before binPMF analysis, but the contamination factor was still resolved. This factor is
318 discussed in more detail in sections 4.1 and 4.4.



319

320 Figure 2 Four-factor result for Range 1, for (a) factor spectral profiles, (b) averaged factor contribution during
 321 the campaign, (c) time series and (d) diurnal trend. Details on the factors' naming schemes are shown in Table
 322 1.

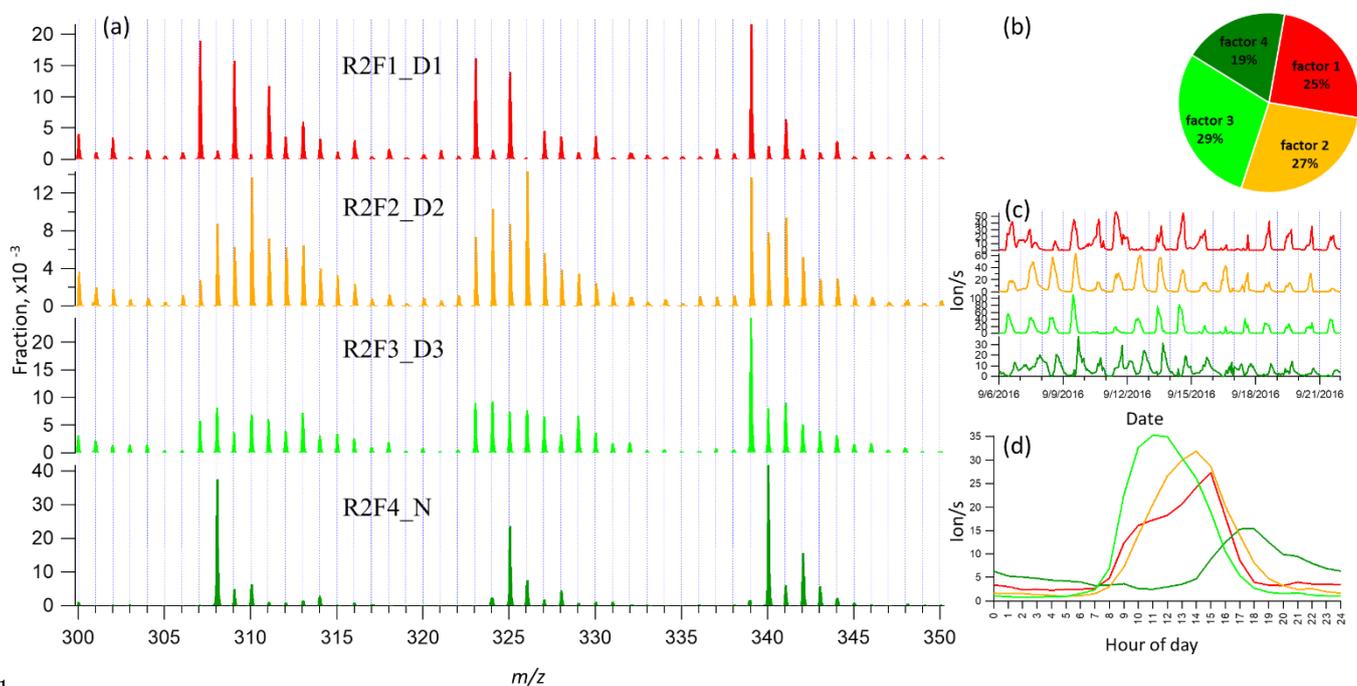
323

324 3.3 binPMF on Range 2 (300-350 Th)

325 This range covers the monoterpene HOM monomer range, and binPMF results have already been
 326 discussed by Zhang et al. (2019) as a first example of the application of binPMF on ambient data.
 327 Our input data here is slightly different. In the previous study, the 10 min automatic zeroing every
 328 three hours was not removed before averaging to 1 hour time resolution while here, we have removed
 329 this data. Overall, the results are similar as in our earlier study, and therefore the result are just briefly
 330 summarized below for further comparison and discussion in Section 4. Similar to Range 1, both the
 331 Q/Q_{exp} (2.2 to 0.6) and unexplained variation (16% to 8%) declined with the increased factor number
 332 from two to ten.

333 When the factor number was two, one daytime and one nighttime factor were separated, with diurnal
 334 peak times at 14:00 and 17:00, respectively. The nighttime factor was characterized by masses at 340
 335 Th, 308 Th and 325 Th (monoterpene ozonolysis HOM monomers (Ehn et al., 2014)) and remained
 336 stable throughout the factor evolution from two to ten factors. With the addition of more factors, no
 337 more nighttime factors got separated while the daytime factor was further separated and more daytime
 338 factors appeared, peaking at various times in the morning (10:00 am), at noon or in the early afternoon
 339 (around 14:00 pm and 15:00 pm). High contribution of 339 Th can be found in all the daytime factor

340 profiles. As the factor number reached six, a contamination factor appeared, characterized by large
 341 signals at 339 Th and 324 Th, showing negative mass defects (Figure S2 in the Supplement). The
 342 factor profile is nearly identical to the contamination factor determined in Zhang et al. (2019), where
 343 the zeroing periods were not removed, causing larger signals for the contaminants. In our dataset,
 344 where the zeroing periods were removed, no sawtooth pattern was discernible in the diurnal trend,
 345 yet it could still be separated even though it only contributed 3% to Range 2. More about the
 346 contamination factors from different ranges will be discussed in Section 4.4. Since the aim of this
 347 study is mainly to explore the benefits of analyzing different ranges of the mass spectrum, we chose
 348 to show the four-factor result below, to simplify the later discussion and comparison. Figure 3 shows
 349 four-factor result of Range 2, with spectral profile, time series, diurnal cycle and averaged factor
 350 contribution during the campaign.



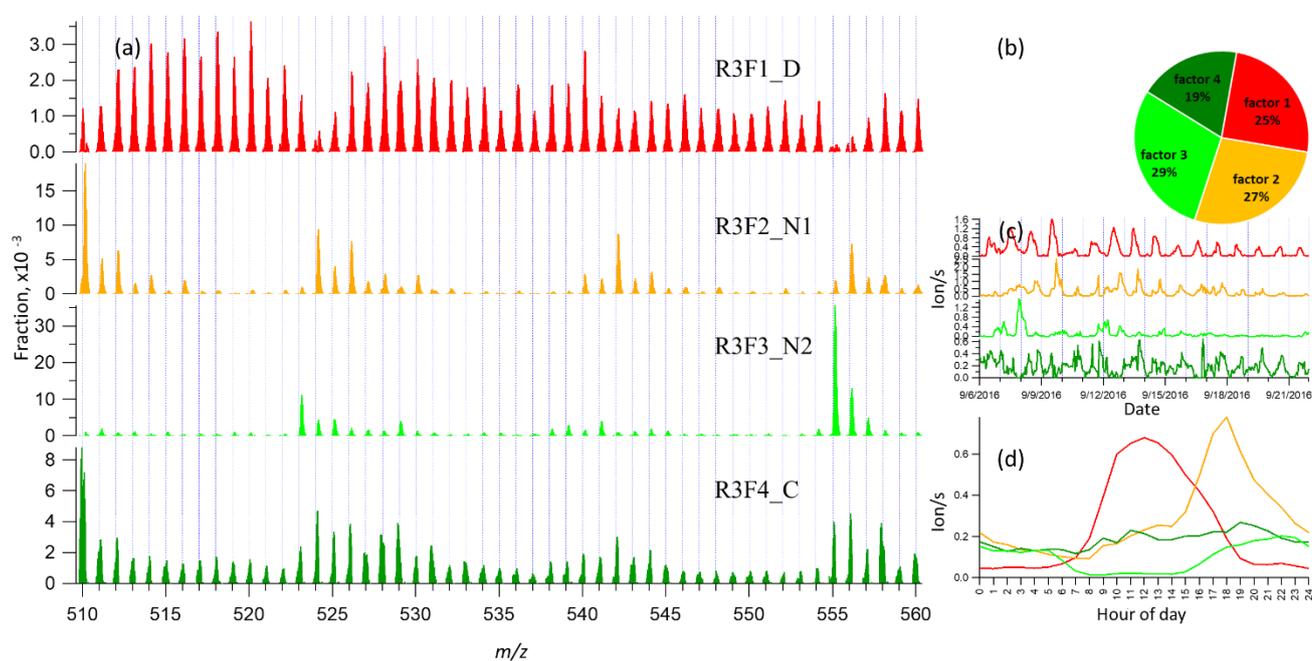
351
 352 Figure 3 Four-factor result for Range 2, for (a) factor spectral profiles, (b) averaged factor contribution during
 353 the campaign, (c) time series and (d) diurnal trend. Details on the factors' naming schemes are shown in Table
 354 1.

355

356 3.4 binPMF on Range 3 (510-560 Th)

357 Range 3 represents mainly the monoterpene HOM dimers (Ehn et al., 2014). Similar to Range 1 and
 358 2, both the Q/Q_{exp} (1.5 to 0.6) and unexplained variation (18% to 15%) showed decreasing trend with
 359 the increased factor number (2-10). As can be seen from Figure 1, data in Range 3 had much lower
 360 signals, compared to that of the Range 1 and 2, explaining the higher unexplained variation for Range
 361 3.

362 In the two-factor result for Range 3, one daytime and one nighttime factor appeared, with diurnal
 363 peak times at noon and 18:00, respectively. The nighttime factor was characterized by ions at 510 Th,
 364 524 Th, 526 Th, 542 Th, and 555 Th, 556 Th, while the daytime factor showed no dominant marker
 365 masses, yet with relatively high signals at 516 Th, 518 Th and 520 Th. As the number of factors
 366 increased to three, one factor with almost flat diurnal trend was separated, with dominant masses of
 367 510 Th, 529 Th, 558 Th. Most peaks in this factor had negative mass defects, and this factor was
 368 again linked to a contamination factor. The four-factor result resolved another nighttime factor with
 369 a dominant peak at 555 Th, and effectively zero contribution during daytime. As the factor number
 370 was further increased, the new factors seemed like splits from previous factors with similar spectral
 371 profiles. We therefore chose four-factor result also for Range 3 (results shown in Fig. 4) for further
 372 discussion.

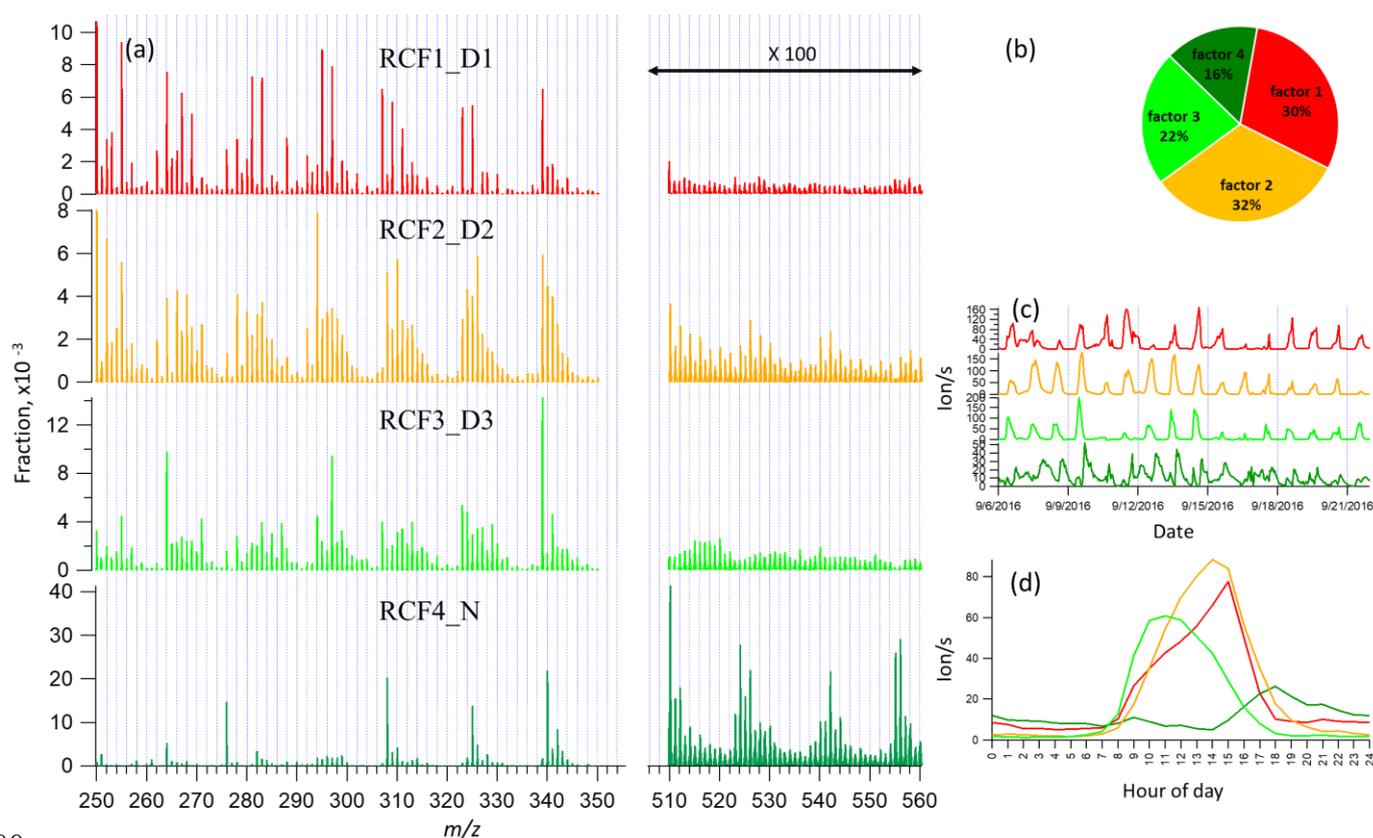


373
 374 Figure 4 Four-factor result for Range 3, for (a) factor spectral profiles, (b) averaged factor contribution during
 375 the campaign, (c) time series and (d) diurnal trend. Details on the factors' naming schemes are shown in Table
 376 1.

377
 378 **3.5 binPMF on Range Combined (250-350 Th & 510-560 Th)**

379 As comparison to the previous three ranges, we conducted the binPMF analysis on Range Combined,
 380 which is the combination of the three ranges. The results of this range are fairly similar to those of
 381 Ranges 1 and 2, as could be expected since the signal intensities in these ranges were much higher
 382 than in Range 3. As the number of factors increased (2-10), both the Q/Q_{exp} (1.3 to 0.6) and
 383 unexplained variation (16% to 8%) showed a decreasing trend.

384 In the two-factor result, one daytime factor and one nighttime factor were separated. In the nighttime
 385 factor, most masses were found at even masses, and the fraction of masses in Range 3 was much
 386 higher than that in daytime factor. In contrast, in the daytime factor, most masses were observed at
 387 odd masses and the fraction of signal in Range 3 was much lower. During the day, photochemical
 388 reactions as well as potential emissions increase the concentration of NO, which serves as peroxy
 389 radical (RO₂) terminator and often outcompetes RO₂ cross reactions in which dimers can be formed
 390 (Ehn et al., 2014). Thus, the production of dimers is suppressed during the day, yielding instead a
 391 larger fraction of organic nitrates, as has been shown also previously (Yan et al., 2016).
 392 With the increase of the number of factors, more daytime factors were resolved with different peak
 393 times. When the factor number reached seven, a clear sawtooth-shape diurnal cycle occurred, i.e. the
 394 contamination factor, caused by the zeroing. As more factors were added, no further nighttime factors
 395 were separated, and only more daytime factors appeared. To simplify the discussion and inter-range
 396 comparison, we also here chose the four-factor result for further analysis, as it already provided
 397 enough information for our main goal in this study. Figure 5 shows the four-factor result of Range
 398 Combined, with spectral profile, time series, diurnal cycle and averaged factor contribution during
 399 the campaign. The signals in range of 510-560 Th were enlarged 100-fold to be visible.



400

401 Figure 5 Four-factor result for Range Combined, for (a) factor spectral profiles, (b) averaged factor
 402 contribution during the campaign, (c) time series and (d) diurnal trend. Details on the factors' naming
 403 schemes are shown in Table 1.

404

405 **4 Discussion**

406 In Section 3, results by binPMF analysis were shown for Ranges 1, 2, 3 and Combined. In this section,
 407 we discuss and compare the results from the different ranges. To simplify the inter-range comparison,
 408 we chose four-factor results for all the four ranges, with the abbreviations shown in Table 1. From
 409 Range 1, three daytime factors and a contaminations factor were separated. In Range 2, three daytime
 410 factors and one nighttime factor (abbreviated as R2F4_N) were resolved. The R2F4_N factor was
 411 characterized by signals at 308 Th ($C_{10}H_{14}O_7 \cdot NO_3^-$), 325 Th ($C_{10}H_{15}O_8 \cdot NO_3^-$), and 340 Th
 412 ($C_{10}H_{14}O_9 \cdot NO_3^-$), and can be confirmed as monoterpene ozonolysis products (Ehn et al., 2014; Yan et
 413 al., 2016). With the increase of factor number to six, the contamination factor got separated also in
 414 this mass range. In Range 3, one daytime factor, two nighttime factors and a contamination factor
 415 were separated. The first nighttime factor (R3F2_N1) had large peaks at 510 Th ($C_{20}H_{32}O_{11} \cdot NO_3^-$)
 416 and 556 Th ($C_{20}H_{30}O_{14} \cdot NO_3^-$), dimer products that have been identified during chamber studies of
 417 monoterpene ozonolysis (Ehn et al., 2014). The molecule observed at 510 Th has 32 H-atoms,
 418 suggesting that one of the RO_2 involved would have been initiated by OH, which is formed during
 419 the ozonolysis of alkenes such as monoterpenes at nighttime (Atkinson et al., 1992; Paulson and
 420 Orlando, 1996). The other nighttime factor (R3F3_N2) was dominated by ions at 523 Th
 421 ($C_{20}H_{31}O_8NO_3 \cdot NO_3^-$) and 555 Th ($C_{20}H_{31}O_{10}NO_3 \cdot NO_3^-$), representing nighttime monoterpene
 422 oxidation involving NO_3 . As these dimers contain only one N-atom, and 31 H-atoms, we can assume
 423 that they are formed from reactions between an RO_2 formed from NO_3 oxidation and another RO_2
 424 formed by ozone oxidation. These results match well with the profiles in a previous study by Yan et
 425 al. (2016). The results of Range Combined are very similar to Range 2, with one nighttime factor and
 426 three daytime factors. The contamination factor was separated with increase of factor number to seven.

427

428

Table 1. Summary of PMF results for the different mass ranges

Range	Factor number	Factor name ^a	Dominant peaks	Peak time
1 (250 - 300 Th)	1	R1F1_D1	250, 255, 295, 297	15:00
	2	R1F2_D2	250, 252, 294	15:00
	3	R1F3_D3	264, 297	11:00
	4	R1F4_C	276	^b
2 (300 - 350 Th)	1	R2F1_D1	307, 309, 323, 325, 339,	15:00
	2	R2F2_D2	310, 326, 339,	14:00
	3	R2F3_D3	339	11:00
	4	R2F4_N	308, 325, 340	18:00

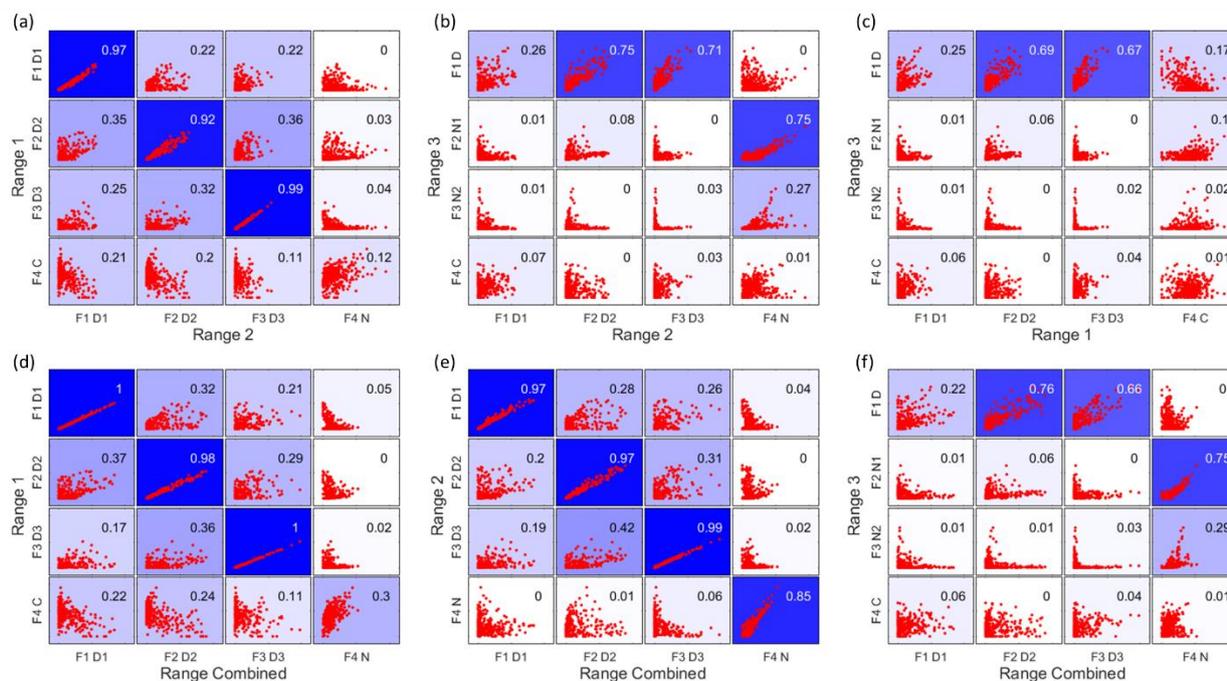
3 (510 – 560 Th)	1	R3F1_D	516, 518, 520, 528, 540	12:00
	2	R3F2_N1	510, 524, 542, 556	18:00
	3	R3F3_N2	523, 555	22:00
	4	R3F4_C	510, 558	- ^b
Combined (1, 2, 3)	1	RCF1_D1	250, 255, 295, 339	15:00
	2	RCF2_D2	250, 252, 294, 339	14:00
	3	RCF3_D3	264, 297, 339	11:00
	4	RCF4_N	308, 340, 510, 524, 555, 556	18:00

429 ^a Factor name is defined with range name, factor number and name. For example, RxFy represents Factor y in Range x.
 430 RC stands for Range Combined. For the factor name, D is short for daytime, N for Nighttime, C for contamination.

431 ^b The contamination factor in Range 1 shows sawtooth pattern; while in Range 3 shows no diurnal pattern.
 432

433 4.1 Time series correlation

434 In Figure 6, the upper panels show the time series correlations among the first three ranges. As
 435 expected based on the results above, generally the daytime factors, and the two nighttime
 436 monoterpene ozonolysis factors (R2F4_N and R3F2_N1) correlated well, respectively. However, the
 437 contamination factors did not show strong correlation between different ranges, even though
 438 undoubtedly from the same source. More about the contamination factors will be discussed in Section
 439 4.1.4. The lower panels in Figure 6 displays the correlations between the first three ranges and the
 440 Range Combined, and clearly demonstrates that the results of Range Combined is mainly controlled
 441 by high signals from Range 1 and 2. More detailed aspects of the comparison between factors in
 442 different ranges is given in the following sections. The good agreements between factors from
 443 different subranges also help to verify the robustness of the solutions.



444 Figure 6 Time series correlations among Range 1, 2, 3 (upper panels a-c), and between the first three ranges
 445 and the Range Combined (lower panels d-f). The abbreviations for different factors are the same in Table 1,
 446 with F for factor, D for daytime, N for nighttime and C for contamination, e.g. F1D1 for Factor 1 daytime 1.
 447

448 The coefficient of determination, R^2 , is marked in each subplot by a number shown in the right upper corners
449 and by the blue colors, with stronger blue indicating higher R^2 .

450

451 **4.2 Daytime processes**

452 **4.2.1 Factor comparison**

453 As mentioned above, with increasing number of factors, usually more daytime factors will be resolved,
454 reflecting the complicated daytime photochemistry. The three daytime factors between Range 1 and
455 2 agreed with each other quite well (Figure 6a). However, R1F1_D1 and R2F1_D1 did not show
456 strong correlation with the only daytime factor in Range 3 (R3F1_D), while the other two daytime
457 factors in both Range 1 and 2, i.e. R1F2_D2, R1F3_D3, R2F2_D2, R2F3_D3, correlated well with
458 R3F1_D from Range 3.

459 The 1st daytime factors from Range 1 and 2, R1F1_D1 and R2F1_D1, were mainly characterized by
460 odd masses 255 Th, 281 Th, 283 Th, 295 Th, 297 Th, 307 Th, 309 Th, 311 Th, 323 Th, 325 Th, 339
461 Th. The factors are dominated by organonitrates. Organic nitrate formation during daytime is
462 generally associated with the termination of RO₂ radicals by NO. This termination step is mutually
463 exclusive with the termination of RO₂ with other RO₂, which can lead to dimer formation. If the NO
464 concentration is the limiting factor for the formation of these factors, the low correlations between
465 the NO-terminated monomer factors, and the dimer factors, is to be expected. In contrast, if the other
466 daytime factors mainly depend on oxidant and monoterpene concentrations, some correlation
467 between those, and the daytime dimer factor, is to be expected, as shown in Figure 6b, c.

468 All the spectral profiles resolved from Range Combined binPMF analysis inevitably contained mass
469 contribution from 510 – 560 Th, even the daytime factor from Range Combined (RCF1_D1) which
470 did not show clear correlation with R3F1_D from Range 3 (Figure 6e).

471 The 2nd and 3rd daytime factors in Range 1 and 2, R1F2_D2, R1F3_D3, R2F2_D2, R2F3_D3, had
472 high correlations with R3F1_D in Range 3. Daytime factors in Range Combined (RCF2_D2 and
473 RCF3_D3) also showed good correlation with R3F1_D in Range 3. However, if we compare R3F1_D
474 and the mass range of 510 – 560 Th of the daytime factors in Range Combined, just with a quick look,
475 we can readily see the difference. The daytime factor separated in Range 3 (R3F1_D) has no obvious
476 markers in the profile. With the increase of factor number (up to ten factors), no clearly new factors
477 were separated in Range 3, but instead the previously separated factors were seen to split into several
478 factors. However, the spectral pattern in R3F1_D is different from that in the mass range of 510 –
479 560 Th in RCF2_D2. The factorization of Range Combined was mainly controlled by low masses
480 due to their high signals. The signals at high masses were forced to be distributed according to the

481 time series determined by small masses. Ultimately, this will lead to failure in factor separation for
 482 this low-signal range.

483 4.2.2 Daytime dimer formation

484 Dimers are primarily produced during nighttime, due to NO suppressing $RO_2 + RO_2$ reactions in
 485 daytime (Ehn et al., 2014; Yan et al., 2016). However, in this study, we found one clear daytime factor
 486 in Range 3 (R3F1_D, peak at local time 12:00, UTC+2) by sub-range analysis. With high loadings
 487 from even masses including 516, 518, 520, 528, 540 Th, this only daytime factor in dimer range
 488 correlated very well with two daytime factors in Ranges 1 and 2 (R1F2_D2, R1F3_D3, R2F2_D2,
 489 R2F3_D3) (Figure 6b and c). Table 2 include the correlation matrix of all PMF and factors and
 490 selected meteorological parameters. Strong correlation between R3F1_D with solar radiation was
 491 found, with $R = 0.79$ (Table 2). This may indicate involvement of OH oxidation in producing this
 492 factor.

493 Table 2 Correlation between factors and meteorological parameters and gases

	R1F1_D1	R1F1_D2	R1F1_D3	R1F1_C	R2F1_D1	R2F2_D2	R2F3_D3	R2F4_N	R3F1_D	R3F2_N1	R3F3_N2	R3F4_C	RCF1_D1	RCF2_D2	RCF3_D3	RCF4_N
O ₃	0.51	0.59	0.35	-0.18	0.47	0.57	0.36	0.43	0.55	0.33	0.27	0.22	0.49	0.57	0.33	0.34
NO	0.13	-0.01	0.24	-0.03	0.18	-0.02	0.24	-0.22	0.13	-0.19	-0.17	0.03	0.13	0.00	0.26	-0.18
NO _x	-0.05	-0.22	-0.10	0.09	-0.01	-0.23	-0.11	-0.13	-0.16	-0.21	-0.04	0.04	-0.04	-0.22	-0.09	-0.11
RH	-0.46	-0.80	-0.63	0.30	-0.43	-0.82	-0.64	-0.27	-0.78	-0.39	-0.07	-0.07	-0.43	-0.82	-0.60	-0.21
T	0.66	0.72	0.40	-0.24	0.65	0.66	0.41	0.39	0.65	0.30	0.14	0.19	0.66	0.68	0.38	0.24
UVB	0.52	0.63	0.82	-0.40	0.52	0.68	0.84	-0.30	0.79	-0.08	-0.27	0.08	0.49	0.68	0.83	-0.29

494
 495 As previous studies have shown, dimers greatly facilitate new particle formation (NPF) (Kirkby et
 496 al., 2016; Troestl et al., 2016; Lehtipalo et al., 2018), and this daytime dimer factor may represent a
 497 source of dimers that would impact the initial stages of NPF in Hyytiälä. Mohr et al. (2017) reported
 498 a clear diel pattern of dimers (sum of about 60 dimeric compounds of $C_{16-20}H_{13-33}O_{6-9}$) during NPF
 499 events in 2013 in Hyytiälä, with minimum at night and maximum after noon, and estimated these
 500 dimers can contribute ~5% of the mass of sub-60 nm particles. The link between the dimers presented
 501 in that paper and those reported here will require further studies, as will the proper quantification of
 502 the dimer factor identified here.

503 4.3 Nighttime processes

504 4.3.1 Factor comparison

505 Since high-mass dimers are more likely to form at night due to photochemical production of NO in
 506 daytime, which inhibits $RO_2 + RO_2$ reactions, Range 3 had the highest fraction of nighttime signals
 507

508 of all the sub-ranges. While Range 3 produced two nighttime factors, Ranges 2 and Combined showed
509 one, and Range 1 had no nighttime factor. The difference between the two results also indicates the
510 advantage of analyzing monomers and dimers separately.

511 The two nighttime factors in Range 3 can be clearly identified as arising from ozonolysis (R3F2_N1)
512 and a mix of ozonolysis and NO₃ oxidation (R3F2_N2) based on the mass spectral profiles, as
513 described above. The organonitrate at 555 Th, C₂₀H₃₁O₁₀NO₃·NO₃⁻, is a typical marker for NO₃
514 radical initiated monoterpene chemistry (Yan et al., 2016). However, several interesting features
515 become evident when comparing to the results of Range 2 and Combined. Firstly, only one nighttime
516 factor (R2F4_N, RCF4_N) was separated in each of these ranges, and that shows clear resemblance
517 with ozonolysis of monoterpenes as measured in numerous studies, e.g. (Ehn et al., 2012;Ehn et al.,
518 2014). Secondly, the high correlation found in Figure 6b between the ozonolysis factors (i.e.,
519 R2F4_N, R3F2_N1, RCF4_N), further supports the assignment. However, factor R2F4_N is the only
520 nighttime factor in the monomer range, suggesting that NO₃ radical chemistry of monoterpenes in
521 Hyytiälä does not form substantial amounts of HOM monomers. The only way for the CI-APi-TOF
522 to detect products of monoterpene-NO₃ radical chemistry may thus be through the dimers, where one
523 highly oxygenated RO₂ radical from ozonolysis reacts with a less oxygenated RO₂ radical from NO₃
524 oxidation.

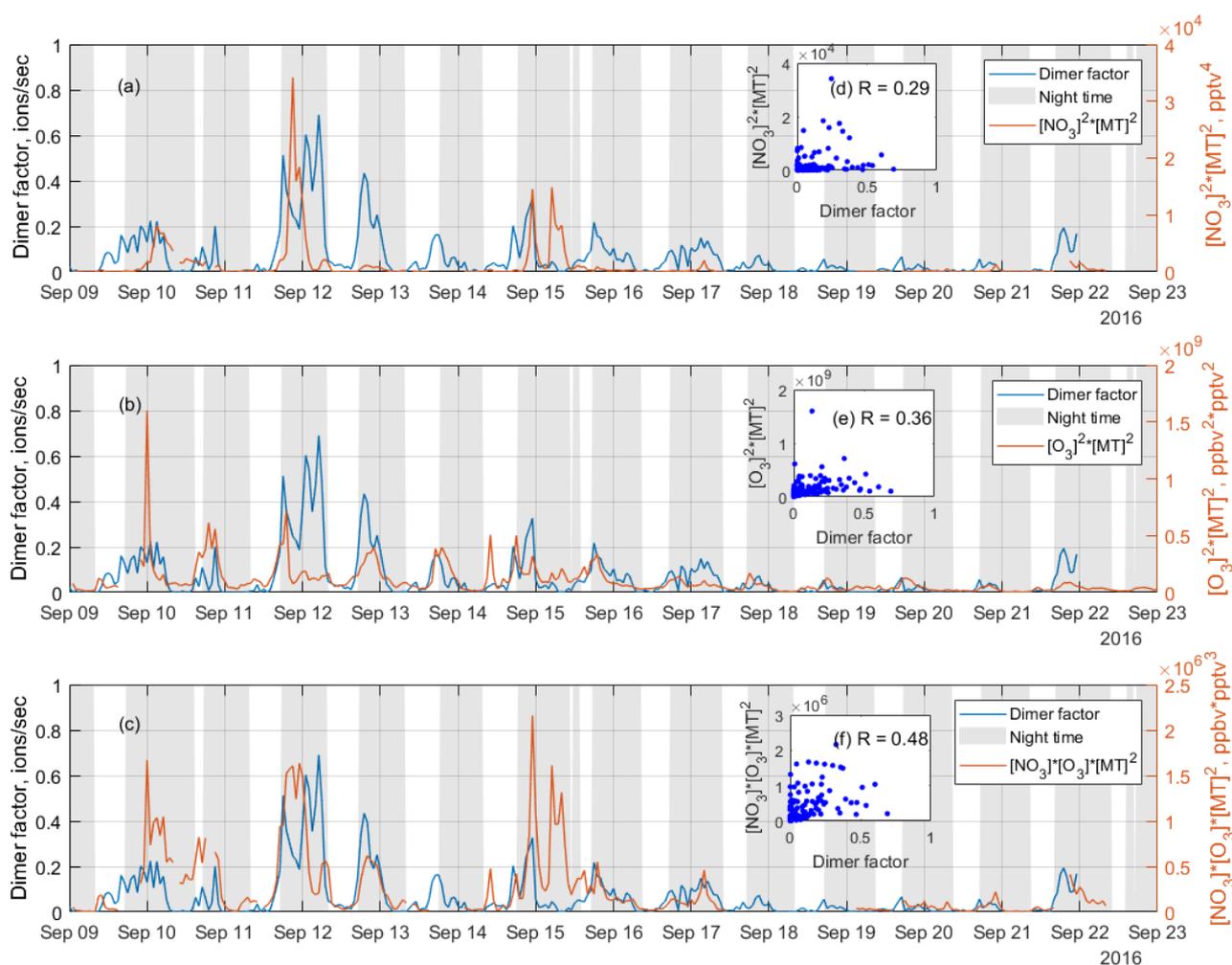
525 In the results by Yan et al. (2016) the combined UMR-PMF of monomers and dimers did yield a
526 considerable amount of compounds in the monomer range also for the NO₃ radical chemistry factor.
527 There may be several reasons for this discrepancy. One major cause for differences between the spring
528 dataset of Yan et al. (2016) and the autumn dataset presented here, is that nighttime concentrations
529 of HOM was greatly reduced during our autumn campaign. The cause may have been fairly frequent
530 fog formation during nights, and also the concentration of e.g. ozone decreased nearly to zero during
531 several nights (Zha et al., 2018). It is also possible that the NO₃ radical-related factor by Yan et al.
532 (2016) is probably a mixture of NO₃ and O₃ radical chemistry, while the monomer may thus be
533 attributed to the O₃ part. Alternatively, the different conditions during the two measurement periods,
534 as well as seasonal difference in monoterpene mixtures (Hakola et al., 2012), caused variations in the
535 oxidation pathways.

536

537 **4.3.2 Dimers initiated by NO₃ radicals**

538 Previous studies show that NO₃ oxidation of α-pinene, the most abundant monoterpene in Hyytiälä
539 (Hakola et al., 2012), produces fairly little SOA mass (yields <4 %), while β-pinene shows yields of
540 up to 53 % (Bonn and Moorgat, 2002;Nah et al., 2016). The NO₃+β-pinene reaction results in low
541 volatile organic nitrate compounds with carboxylic acid, alcohol, and peroxide functional groups (Fry

542 et al., 2014;Boyd et al., 2015), while $\text{NO}_3 + \alpha$ -pinene reaction will typically lose the nitrate functional
 543 group and form oxidation products with high vapor pressures (Spittler et al., 2006;Perraud et al.,
 544 2010). Most monoterpene-derived HOM, including monomers, are low-volatile (Peräkylä et al.,
 545 2020) and thus a low SOA yield indicates a low HOM yield. Thus, while there are to our knowledge
 546 no laboratory studies on HOM formation from NO_3 oxidation of α -pinene, a low yield can be expected
 547 based on SOA studies.



548

549 Figure 7 Time series of the NO_3 oxidation dimer factor (blue line), and the product of (a) $[\text{NO}_3]^2 \times$
 550 $[\text{monoterpene}]^2$, (b) $[\text{O}_3]^2 \times [\text{monoterpene}]^2$, and (c) $[\text{NO}_3] \times [\text{O}_3] \times [\text{monoterpene}]^2$, where $[\]$ represents
 551 concentration in unit of pptv for NO_3 radicals and monoterpene, ppbv for O_3 , while the scatter plots are shown
 552 as inserts, (d), (e), (f), respectively. The scatter plots and correlation coefficients R are only calculated from
 553 nighttime data, which is selected based on solar radiation, to eliminate the influence from daytime oxidation
 554 processes.

555
 556 As discussed above, a dimer factor (R3F2_N2) was identified as being a crossover between NO_3
 557 radical initiated and O_3 initiated RO_2 radicals. Figure 7 shows the time series of this factor, as well as
 558 the product of $[\text{NO}_3]^2 \times [\text{monoterpene}]^2$, $[\text{O}_3]^2 \times [\text{monoterpene}]^2$, and $[\text{NO}_3] \times [\text{O}_3] \times [\text{monoterpene}]^2$.
 559 These products are used to mimic the formation rates of the RO_2 radicals reacting to form the dimers,

560 either from pure NO₃ oxidation (Fig. 7a), pure O₃ oxidation (7b), or the mixed reaction between RO₂
561 from the two oxidants (7c). The NO₃ concentration was estimated in Liebmann et al. (2018) for the
562 same campaign. Monoterpenes were measured using a proton transfer reaction time of flight mass
563 spectrometer (PTR-TOF-MS). More details on measurement of NO₃ proxy and monoterpene can be
564 found in in Liebmann et al. (2018).

565 As shown in Figure 7, the time series of the dimer factor tracks those of [NO₃] × [monoterpene] and
566 [O₃] × [monoterpene] reasonably well, but shows the highest correlation with the product of [NO₃] ×
567 [O₃] × [monoterpene]². This further supports this dimer formation as a mixed processes of ozonolysis
568 and NO₃ oxidation. The heterogeneity of the monoterpene emissions in the forest, and the fact that
569 no dimer loss process is included, partly explain the relatively low correlation coefficients. The
570 sampling inlets for PTR-TOF were about 170 m away from the NO₃ reactivity measurement
571 (Liebmann et al., 2018), which in turn was some tens of meters away from the HOM measurements.
572 Thus, this analysis should be considered qualitative only.

573 The nitrate dimer factor (R3F2_N2) was dominated by the organonitrate at 555 Th,
574 C₂₀H₃₁O₁₀NO₃·NO₃⁻. However, unlike the pure ozonolysis dimer factor which had a corresponding
575 monomer factor (R = 0.86 between factor R2F4_N and R3 F2_N1), this NO₃-related dimer factor did
576 not have an equivalent monomer factor. This suggests that the NO₃ oxidation of the monoterpene
577 mixture in Hyytiälä does not by itself form much HOM, but in the presence of RO₂ from ozonolysis,
578 the RO₂ from NO₃ oxidation can take part in HOM dimer formation. This further implies that,
579 different from previous knowledge based on single-oxidant experiments in chambers, NO₃ oxidation
580 may have a larger impact on SOA formation in the atmosphere where different oxidants exist
581 concurrently. This highlights the need for future laboratory studies to consider systems with multiple
582 oxidants during monoterpene oxidation experiments, to truly understand the role and contribution of
583 different oxidants, and NO₃ in particular.

584 **4.4 Fluorinated compounds**

585 During the campaign, an automated instrument zeroing every three hours was conducted. While the
586 zeroing successfully removed the low-volatile HOM and H₂SO₄, the process also introduced
587 contaminants into the inlet lines, e.g. perfluorinated organic acids from Teflon tubing. Each zeroing
588 process lasted for 10 min. In the data analysis, we removed all the 10-min zeroing periods, and
589 averaged the data to 1-h time resolution, but contaminants were still identified in all ranges by
590 binPMF. However, the correlation between contamination factors from different ranges is low (Figure
591 6c).

592 To future investigate the low factor correlations of the same source, three fluorinated compounds with
593 different volatilities, $(\text{CF}_2)_3\text{CO}_2\text{HF}\cdot\text{NO}_3^-$ (275.9748 Th), $(\text{CF}_2)_5\text{C}_2\text{O}_4\text{H}$ (338.9721 Th), and
594 $(\text{CF}_2)_6\text{CO}_2\text{HF}\cdot\text{NO}_3^-$ (425.9653 Th), were examined in fine time resolution, i.e. 1 min. The time series
595 and 3-h cycle of the three fluorinated compounds were shown in Figure S3 and S4 in Supplement.
596 The correlation coefficients dropped greatly before and after the zero period was removed, from 0.9
597 to 0.3 for R^2 between 276 Th and 339 Th, and 0.8 to 0.1 between 276 Th and 426 Th (Fig. S5a, b).
598 Similar effect is also found with the 1 h averaged data (Fig. S5c, d). It is evident that the three
599 fluorinated compounds were from the same source (zeroing process), but due to their different
600 volatilities, they were lost at different rates. This, in turn, means that the spectral signature of this
601 source will change as a function of time, at odds with one of the basic assumptions of PMF.
602 The analysis of the fluorinated compounds in our system was here merely used as an example to show
603 that volatility can impact source profiles over time. In Figure S5, it can be clearly seen that the profile
604 of Range Combined is noisier than that of Range 3, probably due to the varied fractional contributions
605 of contamination compounds to the profile. In ambient data, products from different sources can have
606 undergone atmospheric processing, altering the product distribution. This analysis highlighted the
607 importance of differences in the sink terms due to different volatilities of the products. This may be
608 an important issue for gas phase mass spectrometry analysis, potentially underestimated by many
609 PMF users, as it is likely only a minor issue for aerosol data, for which PMF has been applied much
610 more routinely. If failing to achieve physically meaningful factors using PMF on gas phase mass
611 spectra, our recommendation is to try applying PMF to sub-ranges of the spectrum, where IVOC,
612 SVOC and (E)LVOC could be analyzed separately.

613

614 **4.5 Atmospheric insights**

615 Based on the new data analysis technique binPMF applied in sub-ranges of mass spectra, we were
616 able to separate two particularly intriguing atmospheric processes, the formation of daytime dimers
617 as well as dimer formation involving NO_3 radicals, which otherwise could not have been identified
618 in our study.

619 With a diurnal peak around noon time, the daytime dimers identified in this study correlate very well
620 with daytime factors in monomer range. Strong correlation between this factor and solar radiation
621 indicate the potential role of OH oxidation in the formation of daytime dimers. By now, very few
622 studies have reported the observations of daytime dimers. As dimers are shown to be able to take part
623 in new particle formation (NPF) (Kirkby et al., 2016), this daytime dimer may contribute to the early
624 stages of NPF in the boreal forest.

625 The second process identified in our study is the formation of dimers that are a crossover between
626 NO_3 and O_3 oxidation. Such dimers have been identified before (Yan et al., 2016). However, we were
627 not able to identify corresponding HOM monomer compounds. This finding indicates that while NO_3
628 oxidation of the monoterpenes in Hyytiälä may not undergo autoxidation to form HOM by themselves,
629 they can contribute to HOM dimers when the NO_3 -derived RO_2 react with highly oxygenated RO_2
630 from other oxidants. Multi-oxidant systems should be taken into consideration in future experimental
631 studies on monoterpene oxidation processes.

632

633 **5 Conclusions**

634 The recent development in mass spectrometry, combined with factor analysis such as PMF, has
635 greatly improved our understanding of complicated atmospheric processes and sources. However,
636 one of PMF's basic assumptions is that factor profiles remain constant in time, yet for atmospheric
637 gas-phase species, reactions and sinks may violate this assumption. In this study, we conducted
638 separate binPMF analysis on three different sub-ranges to explore the potential benefits of such an
639 approach for producing more physically meaningful factors.

640 With binPMF applied on sub-ranges, our study identified daytime dimers, presumably initiated by
641 OH/O_3 with a diurnal peak at around noon, which may contribute to NPF in Hyytiälä. Also, based on
642 the sub-range binPMF analysis, we successfully separated NO_3 -related dimers which did not have a
643 corresponding monomer factor. The NO_3 -related factor was consistent with earlier observations (Yan
644 et al., 2016), but would not have been identified from this dataset without utilizing the different sub-
645 ranges. In future laboratory experiments, more complex oxidation systems may be useful in order to
646 understand the role of NO_3 oxidation in SOA formation. Apart from these two findings, we also find
647 other benefits by applying binPMF on sub-ranges of the mass spectra.

648 First, volatility affects the PMF results. Different compounds emitted from the same source showed
649 different temporal trends, likely due to differences in volatilities. This increased the difficulties for
650 PMF to separate this source in the combined data set, and the resolved profile was less accurate than
651 that of the sub-ranges. Future studies of gas-phase mass spectra should pay attention to this volatility
652 effect on factor analysis.

653 Secondly, chemistry or sources contributing to the particular range can be better separated. Only the
654 binPMF analysis on Range 3, where HOM dimers are typically observed, resolved two nighttime
655 factors, characterized by monoterpene oxidation related to NO_3 and O_3 oxidation.

656 Thirdly, peaks with smaller signal intensities can be correctly assigned. The signal intensities between
657 different parts of the mass spectrum may vary by orders of magnitude. In the combined case, the
658 results were almost completely controlled by the higher signals from smaller masses. The separate

659 analysis on Range 3 allowed the low signals to provide important information. In addition, running
660 binPMF on different separate mass ranges also allows us to compare the factors obtained from the
661 different ranges and help to verify the results.

662

668 **Data availability.** The data used in this study are available from the first author upon request: please
669 contact Yanjun Zhang (yanjun.zhang@helsinki.fi).

677 **Author contribution.** ME and YZ designed the study. QZ and MR collected the data; data analysis
678 and manuscript writing were done by YZ. All coauthors discussed the results and commented the
679 manuscript.

680 **Competing interests.** The authors declare that they have no conflict of interest

681 **Acknowledgements.** We thank the tofTools team for providing tools for mass spectrometry data
682 analysis. The personnel of the Hyytiälä forestry field station are acknowledged for help during field
683 measurements.

684 **Financial support.** This research was supported by the European Research Council (Grant 638703-
685 COALA), the Academy of Finland (grants 317380 and 320094), and the Vilho, Yrjö and Kalle
686 Väisälä Foundation. K.R.D. acknowledges support by the Swiss National Science postdoc mobility
687 grant P2EZP2_181599.

688

689 Reference

- 690 Äijälä, M., Heikkinen, L., Fröhlich, R., Canonaco, F., Prévôt, A. S. H., Junninen, H., Petäjä, T., Kulmala, M.,
691 Worsnop, D., and Ehn, M.: Resolving anthropogenic aerosol pollution types – deconvolution and
692 exploratory classification of pollution events, *Atmos. Chem. Phys.*, 17, 3165-3197, 10.5194/acp-17-
693 3165-2017, 2017.
- 694 Allan, J. D., Jimenez, J. L., Williams, P. I., Alfarra, M. R., Bower, K. N., Jayne, J. T., Coe, H., and Worsnop,
695 D. R.: Quantitative sampling using an Aerodyne aerosol mass spectrometer 1. Techniques of data
696 interpretation and error analysis, *Journal of Geophysical Research: Atmospheres*, 108, 2003.
- 697 Atkinson, R., Aschmann, S. M., Arey, J., and Shorees, B.: Formation of OH radicals in the gas phase reactions
698 of O₃ with a series of terpenes, 97, 6065-6073, 10.1029/92jd00062, 1992.
- 699 Berndt, T., Mentler, B., Scholz, W., Fischer, L., Herrmann, H., Kulmala, M., and Hansel, A.: Accretion Product
700 Formation from Ozonolysis and OH Radical Reaction of α -Pinene: Mechanistic Insight and the
701 Influence of Isoprene and Ethylene, *Environmental Science & Technology*, 52, 11069-11077,
702 10.1021/acs.est.8b02210, 2018a.
- 703 Berndt, T., Scholz, W., Mentler, B., Fischer, L., Herrmann, H., Kulmala, M., and Hansel, A.: Accretion Product
704 Formation from Self- and Cross-Reactions of RO₂ Radicals in the Atmosphere, *Angewandte Chemie
705 International Edition in English* 57, 3820-3824, 10.1002/anie.201710989, 2018b.
- 706 Bertram, T. H., Kimmel, J. R., Crisp, T. A., Ryder, O. S., Yatavelli, R. L. N., Thornton, J. A., Cubison, M. J.,
707 Gonin, M., and Worsnop, D. R.: A field-deployable, chemical ionization time-of-flight mass
708 spectrometer, *Atmospheric Measurement Techniques*, 4, 1471-1479, 10.5194/amt-4-1471-2011, 2011.
- 709 Bianchi, F., Kurtén, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crounse, J. D., Wennberg,
710 P. O., Mentel, T. F., Wildt, J., Junninen, H., Jokinen, T., Kulmala, M., Worsnop, D. R., Thornton, J.
711 A., Donahue, N., Kjaergaard, H. G., and Ehn, M.: Highly Oxygenated Organic Molecules (HOM)
712 from Gas-Phase Autoxidation Involving Peroxy Radicals: A Key Contributor to Atmospheric Aerosol,
713 *Chemical Reviews*, 119, 3472-3509, 10.1021/acs.chemrev.8b00395, 2019.

714 Bonn, B., and Moorgat, G. K.: New particle formation during α - and β -pinene oxidation by O₃, OH and NO₃,
715 and the influence of water vapour: particle size distribution studies, *Atmos. Chem. Phys.*, 2, 183-196,
716 10.5194/acp-2-183-2002, 2002.

717 Boyd, C. M., Sanchez, J., Xu, L., Eugene, A. J., Nah, T., Tuet, W. Y., Guzman, M. I., and Ng, N. L.: Secondary
718 organic aerosol formation from the β -pinene+NO₃ system: effect of humidity and peroxy radical fate,
719 *Atmos. Chem. Phys.*, 15, 7497-7522, 10.5194/acp-15-7497-2015, 2015.

720 Buchholz, A., Lambe, A. T., Ylisirniö, A., Li, Z., Tikkanen, O. P., Faiola, C., Kari, E., Hao, L., Luoma, O.,
721 Huang, W., Mohr, C., Worsnop, D. R., Nizkorodov, S. A., Yli-Juuti, T., Schobesberger, S., and
722 Virtanen, A.: Insights into the O₂-dependent mechanisms controlling the
723 evaporation of α -pinene secondary organic aerosol particles, *Atmos. Chem. Phys.*, 19, 4061-4073,
724 10.5194/acp-19-4061-2019, 2019.

725 Canagaratna, M., Jayne, J., Jimenez, J., Allan, J., Alfarra, M., Zhang, Q., Onasch, T., Drewnick, F., Coe, H.,
726 and Middlebrook, A.: Chemical and microphysical characterization of ambient aerosols with the
727 aerodyne aerosol mass spectrometer, *Mass Spectrometry Reviews*, 26, 185-222, 2007.

728 Canonaco, F., Crippa, M., Slowik, J., Baltensperger, U., and Prévôt, A.: SoFi, an IGOR-based interface for the
729 efficient use of the generalized multilinear engine (ME-2) for the source apportionment: ME-2
730 application to aerosol mass spectrometer data, *Atmospheric Measurement Techniques*, 6, 3649, 2013.

731 Craven, J. S., Yee, L. D., Ng, N. L., Canagaratna, M. R., Loza, C. L., Schilling, K. A., Yatavelli, R. L. N.,
732 Thornton, J. A., Ziemann, P. J., Flagan, R. C., and Seinfeld, J. H.: Analysis of secondary organic
733 aerosol formation and aging using positive matrix factorization of high-resolution aerosol mass spectra:
734 application to the dodecane low-NO_x system, *Atmos. Chem. Phys.*, 12, 11795-11817,
735 10.5194/acp-12-11795-2012, 2012.

736 Crippa, M., Canonaco, F., Lanz, V. A., Äijälä, M., Allan, J. D., Carbone, S., Capes, G., Ceburnis, D., Dall'Osto,
737 M., Day, D. A., DeCarlo, P. F., Ehn, M., Eriksson, A., Freney, E., Hildebrandt Ruiz, L., Hillamo, R.,
738 Jimenez, J. L., Junninen, H., Kiendler-Scharr, A., Kortelainen, A. M., Kulmala, M., Laaksonen, A.,
739 Mensah, A. A., Mohr, C., Nemitz, E., O'Dowd, C., Ovadnevaite, J., Pandis, S. N., Petäjä, T., Poulain,
740 L., Saarikoski, S., Sellegri, K., Swietlicki, E., Tiitta, P., Worsnop, D. R., Baltensperger, U., and Prévôt,
741 A. S. H.: Organic aerosol components derived from 25 AMS data sets across Europe using a consistent
742 ME-2 based source apportionment approach, *Atmos. Chem. Phys.*, 14, 6159-6176, 10.5194/acp-14-
743 6159-2014, 2014.

744 Ehn, M., Kleist, E., Junninen, H., Petäjä, T., Lönn, G., Schobesberger, S., Dal Maso, M., Trimborn, A.,
745 Kulmala, M., Worsnop, D. R., Wahner, A., Wildt, J., and Mentel, T. F.: Gas phase formation of
746 extremely oxidized pinene reaction products in chamber and ambient air, *Atmos. Chem. Phys.*, 12,
747 5113-5127, 10.5194/acp-12-5113-2012, 2012.

748 Ehn, M., Thornton, J. A., Kleist, E., Sipila, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann,
749 R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., Jokinen, T., Schobesberger, S.,
750 Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurten, T., Nielsen, L. B., Jorgensen, S., Kjaergaard,
751 H. G., Canagaratna, M., Dal Maso, M., Berndt, T., Petaja, T., Wahner, A., Kerminen, V.-M., Kulmala,
752 M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic
753 aerosol, *Nature*, 506, 476-479, 10.1038/nature13032, 2014.

754 El Haddad, I., D'Anna, B., Temime-Roussel, B., Nicolas, M., Boreave, A., Favez, O., Voisin, D., Sciare, J.,
755 George, C., Jaffrezo, J. L., Wortham, H., and Marchand, N.: Towards a better understanding of the
756 origins, chemical composition and aging of oxygenated organic aerosols: case study of a
757 Mediterranean industrialized environment, Marseille, *Atmos. Chem. Phys.*, 13, 7875-7894,
758 10.5194/acp-13-7875-2013, 2013.

759 Fry, J. L., Draper, D. C., Barsanti, K. C., Smith, J. N., Ortega, J., Winkler, P. M., Lawler, M. J., Brown, S. S.,
760 Edwards, P. M., Cohen, R. C., and Lee, L.: Secondary Organic Aerosol Formation and Organic Nitrate
761 Yield from NO₃ Oxidation of Biogenic Hydrocarbons, *Environmental Science & Technology*, 48,
762 11944-11953, 10.1021/es502204x, 2014.

763 Guenther, A., Hewitt, C. N., Erickson, D., Fall, R., Geron, C., Graedel, T., Harley, P., Klinger, L., Lerdau, M.,
764 McKay, W. A., Pierce, T., Scholes, B., Steinbrecher, R., Tallamraju, R., Taylor, J., and Zimmerman,
765 P.: A GLOBAL-MODEL OF NATURAL VOLATILE ORGANIC-COMPOUND EMISSIONS,
766 *Journal of Geophysical Research-Atmospheres*, 100, 8873-8892, 10.1029/94jd02950, 1995.

- 767 Hakola, H., Tarvainen, V., Bäck, J., Ranta, H., Bonn, B., Rinne, J., and Kulmala, M.: Seasonal variation of
768 mono- and sesquiterpene emission rates of Scots pine, *Biogeosciences*, 3, 93-101, 10.5194/bg-3-93-
769 2006, 2006.
- 770 Hakola, H., Hellén, H., Hemmilä, M., Rinne, J., and Kulmala, M.: In situ measurements of volatile organic
771 compounds in a boreal forest, *Atmos. Chem. Phys.*, 12, 11665-11678, 10.5194/acp-12-11665-2012,
772 2012.
- 773 Hari, P., and Kulmala, M.: Station for Measuring Ecosystem–Atmosphere Relations (SMEAR II), *Boreal*
774 *Environment Research*, 10, 315-322, 2005.
- 775 Huang, S., Rahn, K. A., and Arimoto, R.: Testing and optimizing two factor-analysis techniques on aerosol at
776 Narragansett, Rhode Island, *Atmospheric Environment*, 33, 2169-2185,
777 [https://doi.org/10.1016/S1352-2310\(98\)00324-0](https://doi.org/10.1016/S1352-2310(98)00324-0), 1999.
- 778 Huffman, J. A., Docherty, K. S., Aiken, A. C., Cubison, M. J., Ulbrich, I. M., DeCarlo, P. F., Sueper, D., Jayne,
779 J. T., Worsnop, D. R., Ziemann, P. J., and Jimenez, J. L.: Chemically-resolved aerosol volatility
780 measurements from two megacity field studies, *Atmos. Chem. Phys.*, 9, 7161-7182, 10.5194/acp-9-
781 7161-2009, 2009.
- 782 Jokinen, T., Sipilä, M., Junninen, H., Ehn, M., Lönn, G., Hakala, J., Petäjä, T., Mauldin Iii, R. L., Kulmala,
783 M., and Worsnop, D. R.: Atmospheric sulphuric acid and neutral cluster measurements using CI-API-
784 TOF, *Atmospheric Chemistry and Physics*, 12, 4117-4125, 10.5194/acp-12-4117-2012, 2012.
- 785 Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M., Yan,
786 C., Almeida, J., Troestl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim, A.,
787 Bernhammer, A.-K., Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X., Craven, J., Dias, A., Ehrhart,
788 S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T., Junninen,
789 H., Kangasluoma, J., Kim, J., Krapf, M., Kuerten, A., Laaksonen, A., Lehtipalo, K., Makhmutov, V.,
790 Mathot, S., Molteni, U., Onnela, A., Peraekylae, O., Piel, F., Petaejae, T., Praplan, A. P., Pringle, K.,
791 Rap, A., Richards, N. A. D., Riipinen, I., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S.,
792 Scott, C. E., Seinfeld, J. H., Sipilae, M., Steiner, G., Stozhkov, Y., Stratmann, F., Tome, A., Virtanen,
793 A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P.,
794 Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R., Baltensperger, U., Kulmala,
795 M., Carslaw, K. S., and Curtius, J.: Ion-induced nucleation of pure biogenic particles, *Nature*, 533,
796 521-526, 10.1038/nature17953, 2016.
- 797 Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T., Petäjä, T., Sipilä,
798 M., Schobesberger, S., Rantala, P., Franchin, A., Jokinen, T., Järvinen, E., Äijälä, M., Kangasluoma,
799 J., Hakala, J., Aalto, P. P., Paasonen, P., Mikkilä, J., Vanhanen, J., Aalto, J., Hakola, H., Makkonen,
800 U., Ruuskanen, T., Mauldin, R. L., Duplissy, J., Vehkamäki, H., Bäck, J., Kortelainen, A., Riipinen,
801 I., Kurtén, T., Johnston, M. V., Smith, J. N., Ehn, M., Mentel, T. F., Lehtinen, K. E. J., Laaksonen, A.,
802 Kerminen, V.-M., and Worsnop, D. R.: Direct Observations of Atmospheric Aerosol Nucleation, 339,
803 943-946, 10.1126/science.1227385 %J Science, 2013.
- 804 Lamarque, J. F., Bond, T. C., Eyring, V., Granier, C., Heil, A., Klimont, Z., Lee, D., Liousse, C., Mievilte, A.,
805 Owen, B., Schultz, M. G., Shindell, D., Smith, S. J., Stehfest, E., Van Aardenne, J., Cooper, O. R.,
806 Kainuma, M., Mahowald, N., McConnell, J. R., Naik, V., Riahi, K., and van Vuuren, D. P.: Historical
807 (1850–2000) gridded anthropogenic and biomass burning emissions of reactive gases and aerosols:
808 methodology and application, *Atmos. Chem. Phys.*, 10, 7017-7039, 10.5194/acp-10-7017-2010, 2010.
- 809 Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Kurtén, T., Worsnop, D. R., and Thornton, J. A.: An Iodide-
810 Adduct High-Resolution Time-of-Flight Chemical-Ionization Mass Spectrometer: Application to
811 Atmospheric Inorganic and Organic Compounds, *Environmental Science & Technology*, 48, 6309-
812 6317, 10.1021/es500362a, 2014.
- 813 Lee, B. H., Lopez-Hilfiker, F. D., D'Ambro, E. L., Zhou, P., Boy, M., Petäjä, T., Hao, L., Virtanen, A., and
814 Thornton, J. A.: Semi-volatile and highly oxygenated gaseous and particulate organic compounds
815 observed above a boreal forest canopy, *Atmos. Chem. Phys.*, 18, 11547-11562, 10.5194/acp-18-
816 11547-2018, 2018.
- 817 Lehtipalo, K., Yan, C., Dada, L., Bianchi, F., Xiao, M., Wagner, R., Stolzenburg, D., Ahonen, L. R., Amorim,
818 A., Baccharini, A., Bauer, P. S., Baumgartner, B., Bergen, A., Bernhammer, A.-K., Breitenlechner, M.,
819 Brilke, S., Buchholz, A., Mazon, S. B., Chen, D., Chen, X., Dias, A., Dommen, J., Draper, D. C.,
820 Duplissy, J., Ehn, M., Finkenzeller, H., Fischer, L., Frege, C., Fuchs, C., Garmash, O., Gordon, H.,
821 Hakala, J., He, X., Heikkinen, L., Heinritzi, M., Helm, J. C., Hofbauer, V., Hoyle, C. R., Jokinen, T.,

822 Kangasluoma, J., Kerminen, V.-M., Kim, C., Kirkby, J., Kontkanen, J., Kürten, A., Lawler, M. J., Mai,
823 H., Mathot, S., Mauldin, R. L., Molteni, U., Nichman, L., Nie, W., Nieminen, T., Ojdanic, A., Onnela,
824 A., Passananti, M., Petäjä, T., Piel, F., Pospisilova, V., Quéléver, L. L. J., Rissanen, M. P., Rose, C.,
825 Sarnela, N., Schallhart, S., Schuchmann, S., Sengupta, K., Simon, M., Sipilä, M., Tauber, C., Tomé,
826 A., Tröstl, J., Väisänen, O., Vogel, A. L., Volkamer, R., Wagner, A. C., Wang, M., Weitz, L., Wimmer,
827 D., Ye, P., Ylisirniö, A., Zha, Q., Carslaw, K. S., Curtius, J., Donahue, N. M., Flagan, R. C., Hansel,
828 A., Riipinen, I., Virtanen, A., Winkler, P. M., Baltensperger, U., Kulmala, M., and Worsnop, D. R.:
829 Multicomponent new particle formation from sulfuric acid, ammonia, and biogenic vapors, 4,
830 eaau5363, 10.1126/sciadv.aau5363 %J Science Advances, 2018.

831 Liebmann, J., Karu, E., Sobanski, N., Schuladen, J., Ehn, M., Schallhart, S., Quéléver, L., Hellen, H., Hakola,
832 H., Hoffmann, T., Williams, J., Fischer, H., Lelieveld, J., and Crowley, J. N.: Direct measurement of
833 NO₃ radical reactivity in a boreal forest, *Atmos. Chem. Phys.*, 18, 3799-3815, 10.5194/acp-18-3799-
834 2018, 2018.

835 Massoli, P., Stark, H., Canagaratna, M. R., Krechmer, J. E., Xu, L., Ng, N. L., Mauldin, R. L., Yan, C., Kimmel,
836 J., Misztal, P. K., Jimenez, J. L., Jayne, J. T., and Worsnop, D. R.: Ambient Measurements of Highly
837 Oxidized Gas-Phase Molecules during the Southern Oxidant and Aerosol Study (SOAS) 2013, *ACS*
838 *Earth and Space Chemistry*, 10.1021/acsearthspacechem.8b00028, 2018.

839 Mohr, C., Lopez-Hilfiker, F. D., Yli-Juuti, T., Heitto, A., Lutz, A., Hallquist, M., D'Ambro, E. L., Rissanen,
840 M. P., Hao, L., Schobesberger, S., Kulmala, M., Mauldin III, R. L., Makkonen, U., Sipilä, M., Petäjä,
841 T., and Thornton, J. A.: Ambient observations of dimers from terpene oxidation in the gas phase:
842 Implications for new particle formation and growth, 44, 2958-2966, 10.1002/2017gl072718, 2017.

843 Nah, T., Sanchez, J., Boyd, C. M., and Ng, N. L.: Photochemical Aging of α -pinene and β -pinene Secondary
844 Organic Aerosol formed from Nitrate Radical Oxidation, *Environmental Science & Technology*, 50,
845 222-231, 10.1021/acs.est.5b04594, 2016.

846 Orlando, J. J., and Tyndall, G. S.: Laboratory studies of organic peroxy radical chemistry: an overview with
847 emphasis on recent issues of atmospheric significance, *J Chemical Society Reviews*, 41, 6294-6317,
848 2012.

849 Paatero, P., and Tapper, U.: Positive matrix factorization: A non-negative factor model with optimal utilization
850 of error estimates of data values, *Environmetrics*, 5, 111-126, 1994.

851 Paatero, P.: Least squares formulation of robust non-negative factor analysis, *Chemometrics and Intelligent*
852 *Laboratory Systems*, 37, 23-35, [https://doi.org/10.1016/S0169-7439\(96\)00044-5](https://doi.org/10.1016/S0169-7439(96)00044-5), 1997.

853 Paatero, P.: The Multilinear Engine—A Table-Driven, Least Squares Program for Solving Multilinear
854 Problems, Including the n-Way Parallel Factor Analysis Model, *Journal of Computational and*
855 *Graphical Statistics*, 8, 854-888, 10.1080/10618600.1999.10474853, 1999.

856 Paciga, A., Karnezi, E., Kostenidou, E., Hildebrandt, L., Psychoudaki, M., Engelhart, G. J., Lee, B. H., Crippa,
857 M., Prévôt, A. S. H., Baltensperger, U., and Pandis, S. N.: Volatility of organic aerosol and its
858 components in the megacity of Paris, *Atmos. Chem. Phys.*, 16, 2013-2023, 10.5194/acp-16-2013-2016,
859 2016.

860 Paulson, S. E., and Orlando, J. J.: The reactions of ozone with alkenes: An important source of HO_x in the
861 boundary layer, 23, 3727-3730, 10.1029/96gl03477, 1996.

862 Peräkylä, O., Riva, M., Heikkinen, L., Quéléver, L., Roldin, P., and Ehn, M.: Experimental investigation into
863 the volatilities of highly oxygenated organic molecules (HOM), *Atmospheric Chemistry and Physics*,
864 20, 649–669, 10.5194/acp-2019-620, 2020.

865 Perraud, V., Bruns, E. A., Ezell, M. J., Johnson, S. N., Greaves, J., and Finlayson-Pitts, B. J.: Identification of
866 Organic Nitrates in the NO₃ Radical Initiated Oxidation of α -Pinene by Atmospheric Pressure
867 Chemical Ionization Mass Spectrometry, *Environmental Science & Technology*, 44, 5887-5893,
868 10.1021/es1005658, 2010.

869 Polissar, A. V., Hopke, P. K., Paatero, P., Malm, W. C., and Sisler, J. F.: Atmospheric aerosol over Alaska: 2.
870 Elemental composition and sources, *Journal of Geophysical Research: Atmospheres*, 103, 19045-
871 19057, 1998.

872 Pope III, C. A., Ezzati, M., and Dockery, D. W.: Fine-particulate air pollution and life expectancy in the United
873 States, *New England Journal of Medicine*, 360, 376-386, 2009.

874 Riva, M., Rantala, P., Krechmer, J. E., Peräkylä, O., Zhang, Y., Heikkinen, L., Garmash, O., Yan, C., Kulmala,
875 M., Worsnop, D., and Ehn, M.: Evaluating the performance of five different chemical ionization

876 techniques for detecting gaseous oxygenated organic species, *Atmospheric Measurement Techniques*,
877 12, 2403-2421, 10.5194/amt-12-2403-2019, 2019.

878 Sekimoto, K., Koss, A. R., Gilman, J. B., Selimovic, V., Coggon, M. M., Zarzana, K. J., Yuan, B., Lerner, B.
879 M., Brown, S. S., Warneke, C., Yokelson, R. J., Roberts, J. M., and de Gouw, J.: High- and low-
880 temperature pyrolysis profiles describe volatile organic compound emissions from western US
881 wildfire fuels, *Atmos. Chem. Phys.*, 18, 9263-9281, 10.5194/acp-18-9263-2018, 2018.

882 Shiraiwa, M., Ueda, K., Pozzer, A., Lammel, G., Kampf, C. J., Fushimi, A., Enami, S., Arangio, A. M.,
883 Fröhlich-Nowoisky, J., Fujitani, Y., Furuyama, A., Lakey, P. S. J., Lelieveld, J., Lucas, K., Morino,
884 Y., Pöschl, U., Takahama, S., Takami, A., Tong, H., Weber, B., Yoshino, A., and Sato, K.: Aerosol
885 Health Effects from Molecular to Global Scales, *Environmental Science & Technology*, 51, 13545-
886 13567, 10.1021/acs.est.7b04417, 2017.

887 Song, Y., Shao, M., Liu, Y., Lu, S., Kuster, W., Goldan, P., and Xie, S.: Source apportionment of ambient
888 volatile organic compounds in Beijing, *Environmental science & technology*, 41, 4348-4353, 2007.

889 Spittler, M., Barnes, I., Bejan, I., Brockmann, K. J., Benter, T., and Wirtz, K.: Reactions of NO₃ radicals with
890 limonene and α -pinene: Product and SOA formation, *Atmospheric Environment*, 40, 116-127,
891 <https://doi.org/10.1016/j.atmosenv.2005.09.093>, 2006.

892 Stocker, T., Qin, D., Plattner, G., Tignor, M., Allen, S., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley,
893 P.: IPCC, 2013: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I
894 to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, 1535 pp, in,
895 Cambridge Univ. Press, Cambridge, UK, and New York, 2013.

896 Troestl, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Ahlm, L., Frege, C., Bianchi, F.,
897 Wagner, R., Simon, M., Lehtipalo, K., Williamson, C., Craven, J. S., Duplissy, J., Adamov, A.,
898 Almeida, J., Bernhammer, A.-K., Breitenlechner, M., Brilke, S., Dias, A., Ehrhart, S., Flagan, R. C.,
899 Franchin, A., Fuchs, C., Guida, R., Gysel, M., Hansel, A., Hoyle, C. R., Jokinen, T., Junninen, H.,
900 Kangasluoma, J., Keskinen, H., Kim, J., Krapf, M., Kuerten, A., Laaksonen, A., Lawler, M., Leiminger,
901 M., Mathot, S., Moehler, O., Nieminen, T., Onnela, A., Petäejae, T., Piel, F. M., Miettinen, P.,
902 Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Sengupta, K., Sipila, M., Smith, J. N.,
903 Steiner, G., Tome, A., Virtanen, A., Wagner, A. C., Weingartner, E., Wimmer, D., Winkler, P. M., Ye,
904 P., Carslaw, K. S., Curtius, J., Dommen, J., Kirkby, J., Kulmala, M., Riipinen, I., Worsnop, D. R.,
905 Donahue, N. M., and Baltensperger, U.: The role of low-volatility organic compounds in initial particle
906 growth in the atmosphere, *Nature*, 533, 527-531, 10.1038/nature18271, 2016.

907 Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.: Interpretation of organic
908 components from Positive Matrix Factorization of aerosol mass spectrometric data, *Atmos. Chem.*
909 *Phys.*, 9, 2891-2918, 10.5194/acp-9-2891-2009, 2009.

910 Yan, C., Nie, W., Aijala, M., Rissanen, M. P., Canagaratna, M. R., Massoli, P., Junninen, H., Jokinen, T.,
911 Sarnela, N., Hame, S. A. K., Schobesberger, S., Canonaco, F., Yao, L., Prevot, A. S. H., Petaja, T.,
912 Kulmala, M., Sipila, M., Worsnop, D. R., and Ehn, M.: Source characterization of highly oxidized
913 multifunctional compounds in a boreal forest environment using positive matrix factorization,
914 *Atmospheric Chemistry and Physics*, 16, 12715-12731, 10.5194/acp-16-12715-2016, 2016.

915 Zha, Q., Yan, C., Junninen, H., Riva, M., Sarnela, N., Aalto, J., Quéléver, L., Schallhart, S., Dada, L.,
916 Heikkinen, L., Peräkylä, O., Zou, J., Rose, C., Wang, Y., Mammarella, I., Katul, G., Vesala, T.,
917 Worsnop, D. R., Kulmala, M., Petäjä, T., Bianchi, F., and Ehn, M.: Vertical characterization of highly
918 oxygenated molecules (HOMs) below and above a boreal forest canopy, *Atmos. Chem. Phys.*, 18,
919 17437-17450, 10.5194/acp-18-17437-2018, 2018.

920 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Ulbrich, I. M., Ng, N. L., Worsnop, D. R., and Sun, Y.:
921 Understanding atmospheric organic aerosols via factor analysis of aerosol mass spectrometry: a review,
922 *Analytical and Bioanalytical Chemistry*, 401, 3045-3067, 10.1007/s00216-011-5355-y, 2011.

923 Zhang, Y., Lin, Y., Cai, J., Liu, Y., Hong, L., Qin, M., Zhao, Y., Ma, J., Wang, X., and Zhu, T.: Atmospheric
924 PAHs in North China: spatial distribution and sources, *Science of the Total Environment*, 565, 994-
925 1000, 2016.

926 Zhang, Y., Cai, J., Wang, S., He, K., and Zheng, M.: Review of receptor-based source apportionment research
927 of fine particulate matter and its challenges in China, *Science of the Total Environment*, 586, 917-929,
928 2017.

929 Zhang, Y., Peräkylä, O., Yan, C., Heikkinen, L., Äijälä, M., Daellenbach, K. R., Zha, Q., Riva, M., Garmash,
930 O., Junninen, H., Paatero, P., Worsnop, D., and Ehn, M.: A novel approach for simple statistical

931 analysis of high-resolution mass spectra, Atmospheric Measurement Techniques, 12, 3761-3776,
932 10.5194/amt-12-3761-2019, 2019.
933