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Diurnal cycle of clouds extending above the tropical tropopause observed by spaceborne lidar

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Abstract

The presence of clouds above the tropopause over tropical convection centers has so far been documented by spaceborne instruments that are either sun-synchronous, or insensitive to thin cloud layers. Here we document, for the first time through direct observation by spaceborne lidar, how the tropical cloud fraction evolves above the tropopause throughout the day. After confirming previous studies that found such clouds are most frequent above convection centers, we show that stratospheric clouds and their vertical extent above the tropopause follow a diurnal rhythm linked to convective activity. The diurnal cycle of the stratospheric clouds displays two maxima: one in the early night (19-20LT) and a later one (00-01LT). Stratospheric clouds extend up to 0.5-1km above the tropopause during nighttime, when they are the most frequent. The frequency and the vertical extent of stratospheric clouds is very limited during daytime, and when present they are found very close to the tropopause. Results are similar over the major convection centers (Africa, South America, Warm pool), with more clouds above land in DJF and less above ocean and JJA.

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1. Scientific context and objectives

The presence of ice clouds near the tropical tropopause has long been documented by in-situ measurements (e.g. Thomas et al., 2002; Jensen et al., 2013; Frey et al., 2014). Detecting occurrences of clouds extending above the tropopause by remote sensing requires documenting the vertical cloud profile with a fine resolution and a high sensitivity to optically thin clouds, which few instruments can reach. Lidar measurements are able to document such occurrences (e.g. Nee et al., 1998), but for a long time were limited to local case studies. Dessler (2009) was the first to use the cloud detections by the CALIPSO lidar (Cloud-Aerosol Lidar Infrared Pathfinder Satellite Observations) to investigate how clouds extend above the tropopause on a global scale. Pan and Munchak (2011) refined the results by using an advanced tropopause dataset. Both studies found that clouds extending into the stratosphere are frequent above seasonal deep convection centers and rarely elsewhere, especially in midlatitudes. Both studies deplored that the fixed overpass local time of the CALIPSO dataset is far from the late afternoon, when land convection is at its maximum. More recently, Wang et al. (2019) documented the presence of laminar cirrus in 10 years of CALIPSO data, and reported a non-negligible cloud amount above the tropopause. Because of the sun-synchronous orbit of CALIPSO, none of these studies were able to document the diurnal cycle of the stratospheric clouds.

Low-stratospheric clouds impact the atmospheric system in several ways. First, their larger heating rate than the clear sky (Corti et al, 2006) increases the upward mass flux and fosters the large-scale upward transport of water above the tropopause. At the hour timescale, the cloud particles penetrating the stratosphere via overshooting convection leads, on the one hand, to a direct stratospheric humidification (Schoeberl et al., 2018; Dauhut et al., 2018). On the other hand, these particles can serve as support for ice-scavenging: under saturated conditions, the water vapor deposits on the particles, which grow and fall out (Corti et al., 2008), decreasing low-stratosphere humidity (Jensen et al., 2013). By all these effects the stratospheric clouds modulate the

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stratospheric water vapor concentrations (Iwasaki et al., 2015) and affect the overall dynamical structure near the tropopause (Corti et al., 2006), at timescales down to one hour. This is why it is important to understand the formation and the sub-daily evolution of such clouds.

Finding the processes responsible for the formation of tropical stratospheric clouds proves difficult, just like with high-tropospheric clouds (Reverdy et al., 2012). Two processes have been mainly proposed. Overshooting convection can lead to the injection of ice crystals into the stratosphere (Dauhut et al., 2018; Lee et al., 2018). Stratospheric cooling triggered by gravity waves (Pfister et al., 2010) could also lead to so-called cloud "in-situ" formation (Pan and Munchak, 2011). The ratio of stratospheric clouds that are formed in-situ has not been estimated yet. The current study does not provide further estimate, but by describing the spatio-temporal evolution of the stratospheric clouds, it highlights how important the convective activity is to drive the stratospheric cloudiness, and how the twice-daily sampling by lidars onboard sun-synchronous platforms can miss the highest and largest stratospheric cloud fraction over certain regions.

In this paper, we document for the first time the diurnal cycle of clouds above the tropopause in the Tropics, and the extent of their penetration in the stratosphere, thanks to the high vertical and temporal resolution of the cloud detection by the CATS (Cloud-Aerosol Transport System) spaceborne lidar (McGill et al., 2015). After describing CATS cloud data, and the method to retrieve the tropopause heights used to detect clouds extending in the stratosphere (Sect. 2), we present maps of stratospheric clouds and document their diurnal cycle in regions of interest (Sect. 3). We then summarise our results and conclude (Sect. 4).

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2. Data and Methods

2.1 CATS Cloud data

The CATS lidar operated from the International Space Station (ISS) between February 2015 and November 2017. It reported profiles at a vertical resolution of 60m every 350m along-track, with an average repeat cycle of nearly 3 days (Yorks et al., 2016). Thanks to the ISS non-synchronous orbit, CATS was able to probe the vertical cloud distribution of a particular region at different times of the day. Aggregating CATS detections over a region of interest and over enough time provides a statistical overview of the diurnal evolution of cloud vertical profiles over that region (Noel et al., 2018).

CATS Level 2 Operational layer files (L2O files, Palm et al., 2016) describe altitudes where cloud layers were detected within profiles of backscatter coefficients measured at 1064nm by the CATS lidar (Pauly et al., 2019), averaged 5km along-track. We considered all such files over the CATS operation period (February 2015 to November 2017) and inspected each 5-km profile within. For profiles located in the Tropics (30S-30N), we inspected each atmospheric layer therein identified as a cloud layer according to the CATS layer type information. As in Noel et al. (2018), we considered layers with a Feature Type Score above 6, to avoid any possibly mislabeled aerosol layers. We flagged the cloud layers with a top altitude above the tropopause. Since any CATS L2O layer entirely above the tropopause is labelled as an aerosol layer (like in CALIPSO, Pan and Munchak, 2011), our study will not include clouds with their base in the stratosphere.

Davis et al (2010) noted that lidars in space may miss the thinnest subvisible cirrus clouds, but with enough spatial averaging optical depths near 0.001 can be detected (Martins et al., 2011). Lidar cloud detections also suffer from a lower sensitivity in the presence of sunlight, which induces significant additional noise in the lidar signal, but climatologies are still relevant (Noel et al., 2018).

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2.2. Tropopause Heights

To obtain the tropopause height, we considered profiles of temperature and pressure from the ERA-5 reanalysis dataset (Albergel et al., 2018). These profiles are available every 6 hours, on 37 vertical levels and a 0.25° x 0.25° horizontal grid. Such profiles in ERA-5 reanalysis agree well with observations in the high tropical troposphere (Podglajen et al. 2014). Using these profiles, we computed the vertical lapse rate profile (as in Reichler et al. 2003), and interpolated it on a 100-m vertical grid. We then applied the WMO criteria defining the presence of a tropopause -- i.e. the lowest altitude at which the lapse rate falls below 2°C/km, provided the lapse-rate between this level and all higher levels within 2 km does not exceed 2°C/km (WMO, 1957). Following the WMO definition, we also allowed for the possibility of a second tropopause if the lapse rate exceeds 3°C/km at least 1 km above the first tropopause. In such a case, we started to look for another tropopause above. To limit computation overhead we constrained the search below 22 km. Using the WMO tropopause definition further allows us to compare our results to previous efforts based on CALIPSO database that used the same definition (Pan and Munchak, 2011).

2.3 Stratospheric cloud detection

For a given CATS 5-km profile (Sect. 2.1), we identified the ERA-5 tropopause height (Sect. 2.2) closest in time and location. Given the 6-hour time resolution of the ERA-5 reanalysis, there is at most 3 hours difference between the observation time and the thermodynamic information used to retrieve the tropopause height. We used the cloud information contained in the 5-km profiles in two ways. First, in 2°x5° lat-lon bins we counted how many profiles contained a cloud extending above the tropopause, compared to the total number of profiles in the bins. Aggregating such numbers observed in JJA and DJF over the CATS operation period produced seasonal maps of above-tropopause cloud amounts (Sect. 3.1). Second, from each CATS 5-km profile we built a

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vertical cloud mask, using the tropopause height as the vertical reference and considering clouds that extend above it. Within regions chosen based on the seasonal maps, we aggregated such cloud masks over the same periods as above, keeping also track of the local time of observation for the considered mask. This produced regional vertical cloud fraction profiles above the tropopause, one profile for each local time of observation (Sect. 3.2).

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3. Results

3.1 Stratospheric cloud distributions

Figure 1 shows the fraction of CATS profiles in which a cloud is detected above the tropopause, in all DJF (top) and JJA (bottom) months of CATS operation.

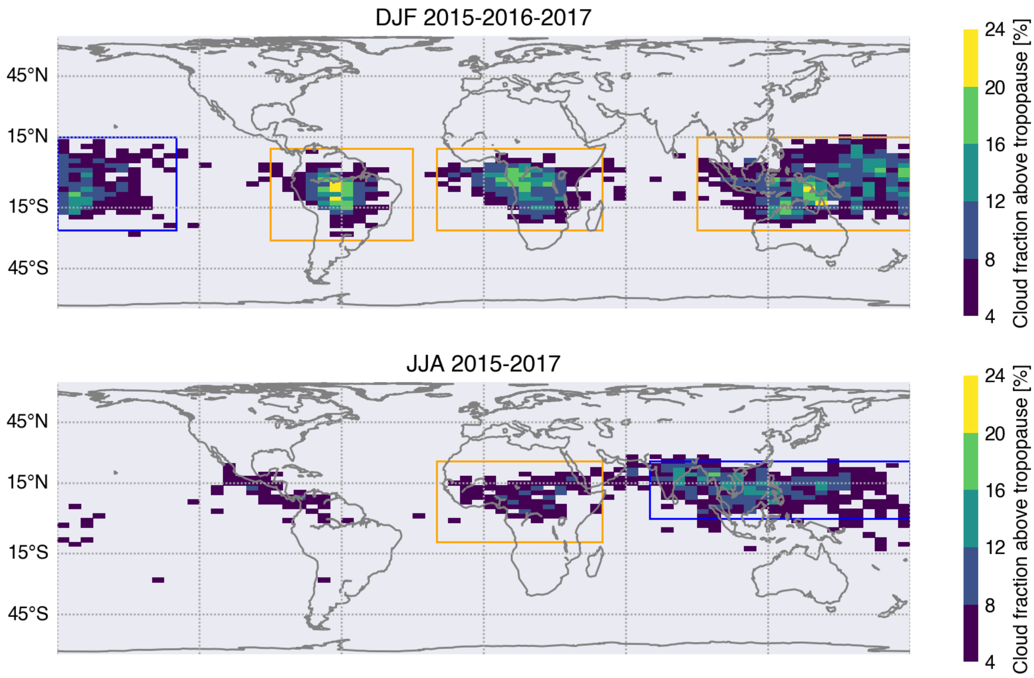

Figure 1*: Tropical low-stratosphere cloud fraction for (top) DJF and (bottom) JJA CATS measurements between Feb 2015 and Nov 2017, calculated by considering all profiles in 2°x5° lat-lon boxes. The rectangles are the regions in which cloud detections are aggregated in the rest of the study. In DJF, from left to right : West Pacific (25S-15N, 180W-130W), South America (30S-10N, 90W-30W), Equatorial Africa (25S-10N, 20W-50E), and South Warm Pool (25S-15N, 90E-180E). In JJA, from left to right : Central Africa (10S-25N, 20W-50E), North Warm Pool (0-25N, 70E-180E). Only detections in the ±30° region are shown here. In the rest of the study, we considered profiles over ocean in blue boxes and profiles over land in orange boxes.*

Figure 1 shows that clouds in the tropical stratosphere are mostly detected over continents (South America, Equatorial Africa and land masses in South Warm Pool in DJF; Central America, Central Africa and land masses in North Warm Pool in JJA). The cloud fraction in the lower stratosphere is

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largest in DJF, up to 24% over central Amazonia and coastal areas in South Warm Pool, and up to 20% over Equatorial Africa. It is significantly lower in JJA, up to 12% over Africa and 16% over the North Warm Pool, even though the lowermost stratosphere (380-420 K potential temperature) is moister in JJA than in DJF (cf. e.g. Fig. 8c in Fueglistaler et al, 2009). Several factors may contribute to this seasonal variation: the density and strength of the convective systems (Liu and Zipser, 2005), their propensity to propagate or to be stationary (Houze et al, 2015), the activity and efficiency of the in situ formation processes (Jensen et al., 2001; Jensen and Pfister, 2004).

The spatio-temporal distribution of the stratospheric clouds is in very good agreement with the 4-year climatology of Pan and Munchak (2011) from CALIPSO observations. The DJF distribution also matches very well the CALIOP cirrus detection at 100 hPa reported by Wang et al (2019) for January 2009. We report though lower cloud frequencies than Wang et al. (2019) which can be explained that we investigate slightly higher levels. Both CATS and CALIPSO datasets find 1) significantly weaker stratospheric cloud fraction in JJA than in DJF, and 2) near-zero stratospheric clouds in the subtropics. These results are also consistent with the CALIPSO cloud fractions near 16km reported by Schoeberl et al. (2019). Since those studies consider cloud detections derived from a spaceborne lidar instrument, over several years for most, their good agreement suggests that the CATS stratospheric cloud detections at 1064 nm are as reliable as the CALIPSO ones at 532 nm. A first conclusion of our results is therefore that CATS measurements strongly support the findings of all other studies using detections of high clouds from CALIPSO data.

Comparing our CATS results to the distributions of clouds at 90 hPa/17 km retrieved from HIRDLS and CALIPSO for 2006-2007 by Massie et al. (2010), we also find good agreement in JJA but larger differences in DJF. In CALIPSO and HIRDLS, the maxima are over the West-to-Central Pacific and the convective spot in South America is shifted West towards the Pacific. This difference can be explained by the annual variability: in DJF 2006-2007 the Southern Oscillation Index indicates

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rather El-Nino conditions, like in DJF 2015-2016 but in contrast with DJF 2016-2017, both being included in our study.

Our results match the overall distributions of cloud tops higher than 17 km retrieved from CloudSat 8-year observations by Kim et al. (2018). In particular, the Warm Pool exhibits the largest area with significant stratospheric cloud fraction, both during DJF and JJA. Little differences between CATS and CloudSat data sets appear in DJF: in the CATS dataset the South America show slightly larger cloud fraction than the Equatorial Africa; in the CloudSat dataset, the largest cloud fractions over Africa are located more south-east (Great Lakes and Madagascar straight). Note that the CloudSat radar samples convection at 1:30 am and pm, potentially missing some continental convective systems. These differences might also be due to the different periods considered: 2006-2014 for CloudSat versus 2015-2017 for CATS. On the contrary, our results contrast with Liu and Zipser (2005) distributions derived from the TRMM Precipitation Radar, where the densities of overshooting systems with tops in the lower stratosphere are remarkably larger in Central America and Central Africa than over the Warm Pool. Since TRMM precipitation radar reflectivities are less sensitive to thin ice particles than CATS and CALIOP lidars, we can interpret this difference by the fact that the American and African systems, though frequently overshooting the stratosphere, produce less stratospheric clouds than the Asian systems.

Finally, our results agree well with the pioneering work of Jensen et al. (1996) who used passive SAGE II observations at 17.5 km for 1989: their cloud fractions are larger because the considered level is closer to the cold point tropopause but the geographical distributions are very close to ours. The differences are: in DJF they observe more clouds over the Atlantic but less over South America, in JJA they observe less clouds over the West Pacific. These differences may be due to the year-to-year variability. The SAGE II instrument relies on a solar occultation method, completely different from the active lidar observation by CATS and CALIPSO.

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3.2 Diurnal cycle of cloud fractions in the tropical stratosphere

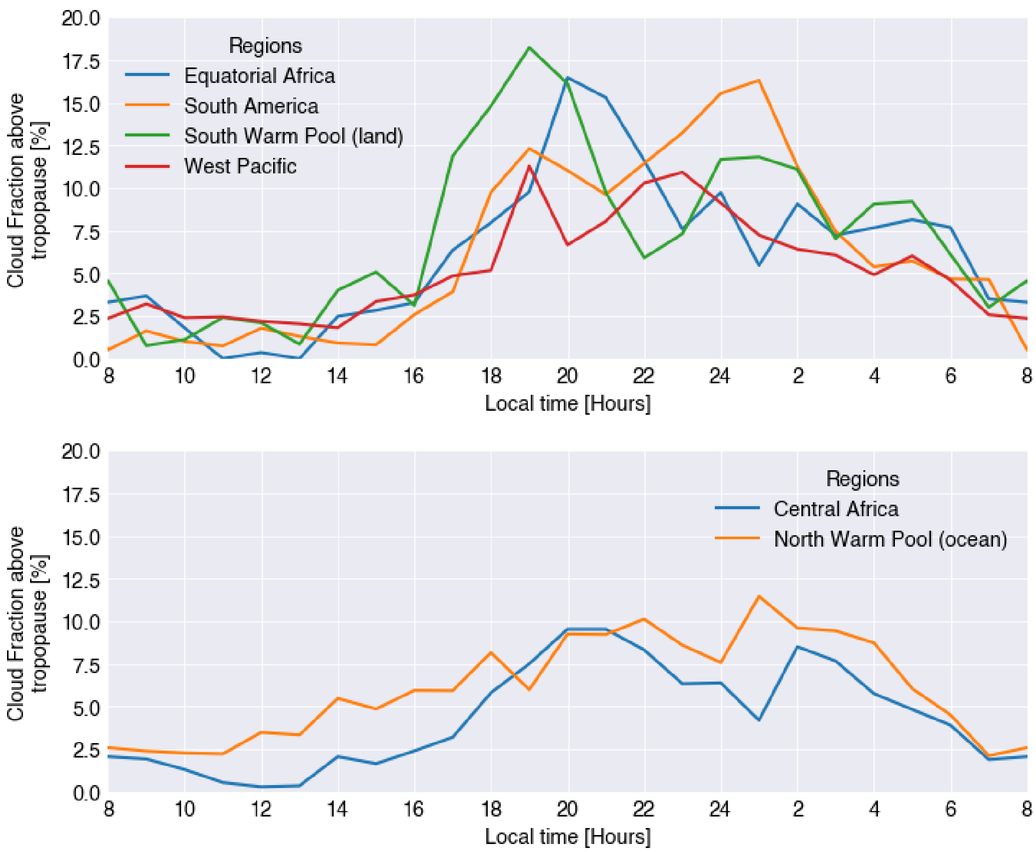

***Figure 2**: Diurnal cycle of stratospheric cloud fraction, by tropical region as in Fig. 1, averaged over DJF (top) and JJA (bottom).*

In contrast with the previous studies, the CATS dataset allows us to analyse the diurnal cycle of the cloud fraction in stratosphere. The cloud fraction at regional scale shows a consistent diurnal cycle, robust over the different regions identified in the previous section (Fig. 2). In particular and in contrast to the diurnal cycle of surface precipitation, there is no land-ocean difference. All exhibit a pronounced minimum about 2-4 % during the day time, from 7 to 16 LT. They all present a first maximum at 19 or 20 LT (early-night peak), up to 18% over South Warm Pool; over North Warm Pool this peak is slightly later. For all regions except South America, this maximum is the largest

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cloud fraction of the day. All regions also present a secondary peak at 0 or 1 LT (2 LT for Central Africa), up to 16% over South America.

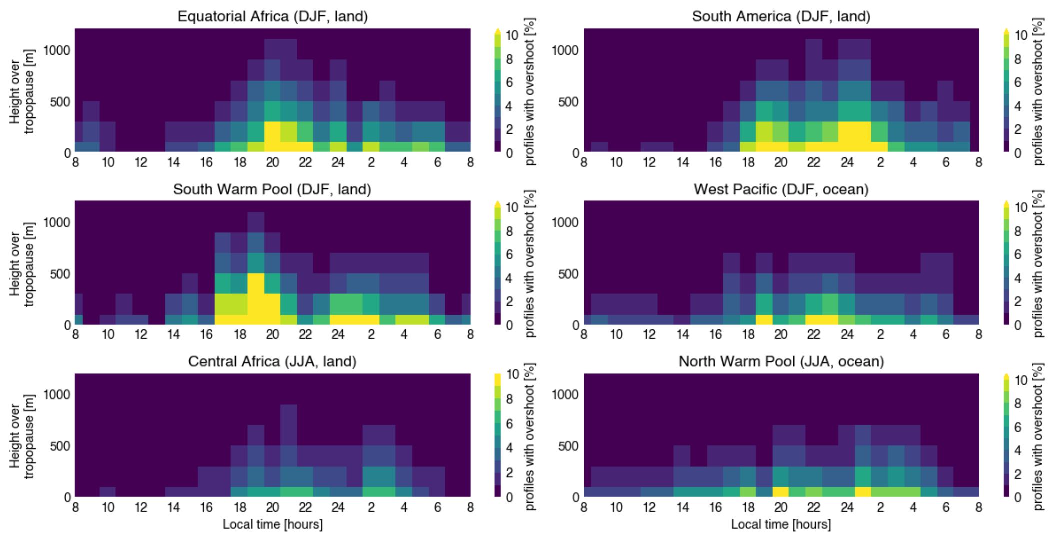

Figure 3: *Diurnal cycle of cloud fraction as a function of height above the tropopause, by tropical region as in Fig. 1, in DJF (top 2 rows) and JJA (bottom row).*

Figure 3 shows how far above the tropopause the clouds extend, depending on the local time in each tropical region (Sect. 3.1). Some regions are considered in DJF, others in JJA, because the stratospheric cloud distribution changes throughout the year (Fig. 1), following the ITCZ position. Patterns appear very consistent in all the regions considered. In all regions the largest cloud fractions are found near the tropopause, with few clouds extending higher. Cloud fractions extend relatively high (up to 1km above the tropopause) during the early night. The first peak of cloud fraction, near 19-20LT (Fig. 2), is associated with the all-day maximum of cloud vertical extent, with clouds in 5% of profiles reaching 1km above the tropopause in DJF regions. During the rest of the night (after 00 LT) clouds are still present but extend less high. During daytime (0600-1800) clouds appear very close to the tropopause. Cloud fractions are overall much smaller in JJA (max 5-10%, bottom row) than in DJF (max 10-12%, rows 1 and 2).

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In addition to describing the evolution of the stratospheric cloud cover at hourly timescales, these observations help interpret observations with limited temporal sampling (Noel et al., 2018). The Microwave Limb Sounder (MLS), like CALIPSO and all other instruments onboard platforms of the A-Train, samples the atmosphere at 01:30 and 13:30 LT, providing one single night and one single day observation. Some authors (e.g., Dion et al., 2019) attempt to retrieve the diurnal cycle of the observed water contents in the tropopause region, combining MLS observations with higher temporal resolution observation of convective activity based on TRMM observation of precipitation. Dion et al. (2019) assumed an in-phase relationship between precipitation and ice water content in the upper troposphere and at the tropopause level. For the stratospheric ice water content, MLS data still provides a too low signal-to-noise ratio. For future investigations, our results indicate that the stratospheric cloud fraction at 13:30 LT is, whatever the region, close to the minimal value of its diurnal cycle, whereas at 01:30 LT it is more typical of the secondary maximum. Carminati et al. (2014) investigated, from MLS measurements between 2005 and 2012, the differences between day and night ice water contents in the upper troposphere and the tropopause level. Unlike the stratospheric cloud fraction, tropopause ice water contents are larger at 13:30 LT than at 01:30 LT over Equatorial Africa during DJF, Central Africa during JJA, and over South America during both seasons. A possible explanation to reconcile our results is that tropopause ice water content is more sensitive to fresh convective activity (very deep convection occurrence) whereas the stratospheric cloud cover is more sensitive to the diffusion of the injected ice in the stratosphere.

4. Conclusion

Our results show how clouds in the tropical stratosphere are strongly concentrated above deep convection centers, are almost absent in subtropical regions, are more frequent in DJF than JJA, and over land than over ocean. In addition to these results, which are consistent with most previous studies, we also show that both the cloud fraction and its extension above the tropopause follow a diurnal rhythm with a maximum during the early nighttime and a near-zero minimum during daytime. During daytime, the stratospheric clouds are limited to the first hundred meters above the tropopause. During nighttime, significant average cloud fraction is found up to 1 km above the tropopause. A secondary maximum of stratospheric cloud fraction is observed over all regions, generally little after midnight. Further investigation is necessary to identify the processes driving this diurnal cycle, and leading in particular to the minimal stratospheric cloud fraction during daytime and the secondary peak during nighttime, both consistent over all regions. Finally further research is needed to understand why the timing of this diurnal cycle is very similar over land and over ocean.

Author contribution. TD and VN designed the data analyses and VN carried them out. TD prepared the manuscript with contributions from VN and ID.

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