Author Comment in response to Referee Comment 2 from Anonymous Referee #1

We thank Anonymous Referee #1 for his/her comments. In the following we answer each of them. The line numbers refers to the version online since 8 October 2019, that differ of 6 with the line references usd by Referee #1. Note that we would like to slightly change the title for: “The diurnal cycle of the clouds extending above the tropical tropopause observed by spaceborne lidar”.
Also, to emphasize our results, a sentence will be added line 236:
“These results highlight how much the evolution of stratospheric clouds can be undersampled by other spatial instruments restricted to 01:30 and 13:30 LT, that then miss for instance the first maximum and the deepest development of stratospheric clouds in the early night.”

- ✔ Referee #1:

The work focus on the use of the CATS lidar to look at stratospheric clouds. It provides a documentation of the diurnal cycle of these clouds over 5 regions of the tropical belt, during 3 summers (JJA) and 2 winters (DJF). Despite the interest of these clouds and the very attractive CATS dataset, I find the paper very light in terms of interpretation.This is currently only providing a documentation already available (as stated by the authors). There is no link with the diurnal cycle of convection (except a very very short mention to surface precipitation) and to the convective overshootings in general, while the regions that are sampled clearly link the observed clouds to the convective activity of the tropics. I thus propose a major revision: as it, the study do not bring somethingreally new, except the view proposed by CATS. This study could be very valuable if the link to convection, and its diurnal cycle, was made.

Authors’ response:
We agree to develop the discussion around the convection’s diurnal cycle, and convection overshoots. The manuscript will be changed as descibed below.

Authors’ changes in manuscript:
We will add the following paragraph after the line 193 and the new paragraph [*] described below:
“The very deep convection transports cloudy air masses beyond the tropopause via overshoots and then directly contributes to the stratospheric cloud fraction (Dauhut et al., 2016 and 2018). The diurnal cycle of the stratospheric cloud fraction observed by CATS can at the first order be explained by the diurnal cycle of very deep convection over land (Liu and Zipser, 2005), especially (i) the minimal value during daytime, and (ii) the first peak in the early evening. This first peak occurs with a delay of 3 to 4 hours compared to the very deep convection maximum. As the dataset used by Liu and Zipser (2005) is more sensitive to overshoots freshly developed into the stratosphere, this delay can be explain by the subsequent horizontal expansion of the overshoots and their spread by the winds (Dauhut et al., 2018; Lee et al., 2019). The convective generation of gravity waves, that produce transcient cooling off the convective centres and in some conditions trigger cloud formation, can also contribute to the increase of the stratospheric cloud fraction after the maximum of the very deep convection, and then explain the delay of the first peak and potentially the second peak. It may also explain the similar diurnal cycle over the ocean regions, either close (South Warm Pool Ocean) or remote (West Pacific) from land masses. This process remains to be investigated.”

The sentences lines 236-238 will be changed for:
“The very convective activity over tropical lands drives most of this diurnal cycle, and leadsing in particular to the minimal stratospheric cloud fraction during daytime and the second peak during nighttime, both consistent over all regions. Further investigation is though necessary to describe how convection contributes to this diurnal cycle, and to assess the role of other processes leading to stratospheric cloud formation like the gravity waves.”

✔ Referee #1:

Other comments - The first section is strangely organized : I suggest to start withthe current 2nd paragraph (starting with "Low-stratospheric clouds impact (...)", endingwith "evolution of such clouds") and then continue with the current 1st paragraph. The idea would be first to introduce the scientific question (low stratospheric clouds & their impact on the atmosphere) and second to present the way it will be looked at (spaceborne lidar).

Authors' response and Authors' changes in manuscript:
We will exchange paragraph one and two (lines 23 to 50), thank you for the suggestion.

✔ Referee #1:

At the end of the first section, the CATS liar is quickly mentioned. Since the CALIOP lidar is mentioned earlier, in the same section, it is not clear why to use CATS rather than CALIOP. The reader has to go to section 2.1 to understand why CATS is used. So a paragraph stating clearly the issue (obviously related to the 1:30pm/1:30am sampling, giving very few information on the diurnal cycle) with CALIOP is missing in section 1.

Authors' response:
We will add a new paragrpah in the section 1 to motivate the use of CATS lidar observation

Authors' changes in manuscript:
The following paragraph will be added before line 51:
"Describing the diurnal evolution of the high-altitude clouds from a global perspective becomes possible with the CATS (Cloud-Aerosol Transport System) lidar operated from the International Space Station (ISS) between February 2015 and November 2017 (McGill et al., 2015). Thanks to the ISS non-synchronous orbit, CATS was able to probe the vertical cloud distribution of a particular region at different times of the day (not only 0130 and 1330 Local Time like CALIPSO instruments). Aggregating CATS detections over a region of interest and over enough time provides a statistical overview of the diurnal evolution of cloud vertical profiles over that region (Noel et al., 2018). Our work aims at using CATS observations to describe and understand better the diurnal evolution of the cloud fraction in the tropical stratosphere."
line 63 the acronym expansion will be deleted
line 64 the reference will be deleted
lines 70-76 will be deleted, only the following sentence is kept: "Between February 2015 and November 2017, the CATS lidar reported profiles at a vertical resolution of 60m every 350m along-track, with an average repeat cycle of nearly 3 days (Yorks et al., 2016)." and merged with the paragraph after.
line 80, the time reference in the brackets will be deleted.

✔ Referee #1:

Section 3 : If I count correctly DJF2015 does not exist since the CATS data start inFeb 2015. So there are 2 DJF and 3 JJA.

Authors' response and Authors' changes in manuscript:
Indeed, only February and December are available in 2015. To make it more accurate we will change the titles of the panels of Figure 1 for "DJF months from Feb. 2015 to Nov. 2017 included" and "JJA months from Feb. 2015 to Nov. 2017 included".

- ✔ Referee #1:

Section 3: the word "level" is used from time to time instead of "altitude" (lines 139;171). Please use "altitude" when it is adapted. The term level is too vague.

Authors' response:
We will change the text accordingly.
Authors' changes in manuscript:
We changed "level" for "altitude in lines: 99 (twice), 145, 177

- ✔ Referee #1:

Section 3 / lines 147 to 175 : 3 paragraphs are dedicated to evaluate the distribu-tions found with CATS with respect to previous works performed with other instruments(HIRDLS - by the way, please expand : the reader don't know this one; CloudSat andSAGE-II). As underlined by the authors, it is difficult to compare the values obtainedwith the mentioned papers since they don't look at the same period. So the year-to-year variability explains largely the differences. That is why I don't understand thestructure of these 3 paragraphs. The year-to-year variability should be written at thebeginning, to explicitly say that the numbers found cannot be compared, and then go to the specificities of each instruments to explain the differences (occultation, radar,etc..). Now it is too repetitive.

Authors' response:
We agree, the three paragraphs will be summarized into one.

Authors' changes in manuscript:
lines 153 to 181 will be replaced by the following paragraph:
"Our CATS results are also in very good agreement with the distributions of clouds near the tropopause from other space instruments: 2006-2007 HIRDLS (High Resolution Dynamics Limb Sounder) reported by Massie et al. (2010), 2006-2014 CloudSat observations (Kim et al., 2018), and the pioneering 1989 passive Stratospheric Aerosol and Gas Experiment (SAGE) II observations (Jensen et al., 1996). Besides the specificity in the cloud detection method employed by each instrument (occultation for HIRDLS and SAGE II, radar backscattering for CloudSat), the little differences between the distributions mostly come from the year-to-year variability. Larger differences can be found with the distributions of clouds penetrating the tropical tropopause derived from the 1998-2000 and 2002-2003 observations by the Tropical Rainfall Measuring Mission (TRMM) Precipitation Radar (Liu and Zipser, 2005). The densities of overshooting systems with tops in the lower stratosphere (on which Liu and Zipser (2005) focused rather than all stratospheric clouds) are remarkably larger in Central America and Central Africa than over the Warm Pool. Since TRMM precipitation radar reflectivities are less sensitive to thin ice particles than CATS and CALISPO lidars, we can interpret this difference by the fact that the American and African systems, though frequently overshooting the stratosphere, produce less thin stratospheric clouds than the Asian systems, or other in-situ processes (like gavity wave cooling) are more efficient to produce stratospheric clouds over Asia than Ameria and Africa.

- ✔ Referee #1:

Also, I wonder if there is no diurnal cycle studies of these clouds performed using ground-based lidar. Is that so ?

Authors' response:
Ground-based lidars have indeed already been used to document the diurnal evolution of high-altitude cirrus clouds. For instance using ground-based lidar measurements: Gouveia et al. (2017) documented the evolution of the integrated cloud fraction (no vertical distribution) over Amazonia,

Sassen et al. (2003) documented the diurnal evolution of the composition of cirrus clouds over Salt Lake City, and Dupont et al. (2010) did the same over the SIRTA observatory in France. However, using ground-based lidar to document optically thin clouds extending above the tropopause, as we did in the present paper, is difficult for two reasons: 1) as our Fig. 1 shows, these clouds occur primarily in regions where operational ground-based sites are absent or very few (Pacific ocean, equatorial Africa, South America), and 2) these clouds are mainly associated with deep convection, which implies the presence of optically thick cloud systems in the troposphere beneath that will make in most cases impossible the successful probing of optically thin clouds near the tropopause due to the attenuation of lidar signal. This explains why we are not aware so far of any article documenting the evolution of such clouds based on ground-based lidars.

Dupont, J.-C., Haeffelin, M., Morille, Y., Noël, V., Keckhut, P., Winker, D., Comstock, J., Chervet, P., and Roblin, A. (2010), Macrophysical and optical properties of midlatitude cirrus clouds from four ground-based lidars and collocated CALIOP observations, J. Geophys. Res., 115, D00H24, doi:10.1029/2009JD011943

Gouveia, D. A., Barja, B., Barbosa, H. M. J., Seifert, P., Baars, H., Pauliquevis, T., and Artaxo, P.: Optical and geometrical properties of cirrus clouds in Amazonia derived from 1 year of ground-based lidar measurements, Atmos. Chem. Phys., 17, 3619–3636, https://doi.org/10.5194/acp-17-3619-2017, 2017

Sassen, K., Liou, K.-N., Takano, Y., and Khvorostyanov, V. I. (2003), Diurnal effects in the composition of cirrus clouds, Geophys. Res. Lett., 30, 1539, doi:10.1029/2003GL017034, 10.

Authors' changes in manuscript:
- line 28: we added a reference to Gouveia et al. (2017)
- The following paragraph will be added after line 50:

"The diurnal evolution of the high-altitude cirrus clouds have been documented over some specific sites using ground-based lidars (Sassen et al. 2003; Dupont et al., 2010; Gouveia et al. 2017). Gouveia et al. (2017) documented the evolution of the integrated cloud fraction (no vertical distribution) over Amazonia, Sassen et al. (2003) documented the diurnal evolution of the composition of cirrus clouds over Salt Lake City, and Dupont et al. (2010) did the same over four observatories in France and in the United-States. However, using ground-based lidar to document optically thin clouds extending above the tropopause is difficult for two reasons: 1) as the studies based on CALIPSO observations show, these clouds occur primarily in regions where operational ground-based sites are absent or very few (Pacific ocean, equatorial Africa, South America), and 2) these clouds are mainly associated with deep convection, which implies the presence of optically thick cloud systems in the troposphere beneath that will make in most cases impossible the successful probing of optically thin clouds near the tropopause due to the attenuation of lidar signal. This explains why the ground-based lidars do not document the diurnal cycle of the stratospheric clouds with a satisfying spatial and temporal coverage."

- [*] The following paragraph will be added directly after line 193:

"The cirrus clouds observed over Amazonia by gound-based lidar (Gouveia et al., 2017) shows a very similar diurnal cycle: a first peak in the early night (at 18-19 LT), a second peak later in the night (at 2-3 LT). Though Gouveia et al. (2017) do not consider the cloud above the tropopause only, their distinction between subvisible, thin and opaque cirrus indicates that the opaque cirrus are predominant during the early night (18-21 LT) and the thin cirrus (and subvisible ones during the dry season) dominate during the later night (from 0 and 2 LT onwards, in wet and dry seasons, respectively)."

- ✔ Referee #1:

Fig 2: please add DJF/JJA at the top of the panels as a quick reminder. Also, the color code for the lines should refer to the color code of Fig 1 : blue for oceanic regions/ land for orange regions. This would make a more logical reading of the figures. Also,I don't see why the day starts at 8am. Is there a particular reason for that ? If there is no specific reason, then it should start at 0:00.

Authors' response:
We will add "DJF months from Feb. 2015 to Nov. 2017 included" and "JJA months from Feb. 2015 to Nov. 2017 included" at the top of the panels and change the colors to have warm color for land and cold for ocean, thank you for the suggestions. We decided to present the diurnal cycle from 8 to 8LT, roughly matching with day time and then night time, to increase the readibility of the figure: the two maxima of the diurnal cycle appear more clearly (they are not separated by the midnight line). See for instance the fig. 4a of Gouveia et al. (2017) to see the difficulty to compare evening and late night when diurnal cycle is presented from 0 to 0 LT.

Authors' changes in manuscript:
The colors and panels' title used in Figure 2 will be changed.

- ✔ Referee #1:

line 186: it is stated that all regions present a secondary peak at 0-1LT. The EquatorialAfrica doesn't have a secondary peak, while for the North Warm Pool, the 1:00LT peak is the first peak in terms of intensity. Please revise.

Authors' response:
Thank you for the comment, we will rephrase these sentences.

Authors' changes in manuscript:
Sentences lines 189-193 will be rephrased with:
"They all present a first maximum at 19 or 20 LT (early-night peak), up to 16.5% over Equatorial Africa. For all regions except South America and North Warm Pool, this maximum is the largest cloud fraction of the day. All regions also present a second peak (late-night peak) at 0 or 1 LT (23 LT for West Pacific and 2 LT for Central Africa), up to 16.5% over South America. The midnight peak over Equatorial Africa is less clear than over the other regions because of the large variations between 23 and 3 LT. The capability of a longer dataset to produce a clearer signal remains to be investigated."
All other occurences of "secondary" will also be changed for "second" : lines 218, 235, 238.

- ✔ Referee #1:

Figure 3: the darker color for the very small values of % makes the reading of thefigure difficult. Please put white when it is 0, so that the small values of % can still be readable. I have the same comment than for fig 2 and the start of the x-axis at 8am.

Authors' response and Authors' changes in manuscript:
We will change the color palette of Figure 3 to have white for zero values. Like for Fig. 2 we decided to present the diurnal cycle from 8 to 8LT to make the two peaks of the diurnal cycle more clear and easily comparable.

Author Comment in response to Referee Comment 1 from Anonymous Referee #2

We thank Anonymous Referee #2 for his/her comments. In the following we answer each of them. The line numbers refers to the version online since 8 October 2019. Note that we would like to slightly change the title for: “The diurnal cycle of the clouds extending above the tropical tropopause observed by spaceborne lidar”.
Also, to emphasize our results, a sentence will be added line 236:
“These results highlight how much the evolution of stratospheric clouds can be undersampled by other spatial instruments restricted to 01:30 and 13:30 LT, that then miss for instance the first maximum and the deepest development of stratospheric clouds in the early night.”

- ✔ Referee #2:

General Impressions:
This study uses the CATS lidar to estimate the stratospheric cloud fraction over tropics. The spatial distribution and the diurnal cycle of the stratospheric cloud fractionare analyzed and documented. The study also discusses the regional and seasonal differences of the stratospheric cloud fraction. The similarities and differences between their results and previous results are discussed as well. As this paper pointed out, it is a rare opportunity to study diurnal cycle of thin clouds based on the vertical cloud profiles. Therefore, I am overall supportive of the study. However, I feel the current manuscript needs a bit more detail (see some suggestions below).

Author Comment:
Indeed the CATS observations provide us an unprecedented opportunity to study the diurnal cycle of the clouds above the tropopause. In the proposed changes, we will add several new paragraphs about: the diurnal cycle of high-altitude cirrus by ground-based lidar, the comparison between our results and ground-based lidar observations, the link between the diurnal cycle of very deep convection and stratosphere cloud faction (see Authors’ changes in Manuscript in the response to Referee Comment 2 from Referee #1).

- ✔ Referee #2:

Specific Suggestions:
Figure 3: Comparisons between DJF and JJA are hard to interpret because the locations are differently selected by different seasons. Are thedifferences we see here due to the seasonal differences or due to the regional differences?

Authors’ response:
The objective is not to compare DJF and JJA but to characterise the cloud fraction in the stratosphere where and when they are the most present, as illustrated in Figure 1. So the differences are firstly due to regional differences. To be more explicit, the lines 196-198 are changed for the text below.

Authors’ changes in manuscript:
"To characterise the diurnal cycle of the cloud extent in the stratosphere, the cloud fraction is represented in terms of height above the tropopause and local time (Fig. 3) over the tropical regions where the cloud fraction is the largest (Sect. 3.1). Each tropical region is considered either in DJF or in JJA to match the season when the stratospheric clouds are the most frequent (Fig. 1)."

- ✔ Referee #2:

The paper concluded that the most of the findings are consistent with previous studies based on satellite observations, which means that the most of the findings are alreadyknown. I understand the uniqueness of this study is to use lidar instruments to understand the cloud fraction. However, the whole story feels a bit thin. Maybe compare theresults with in-situ measurements to add more insight?

Authors' response:
Thank you for the suggestion. However, most of our findings concern the diurnal cycle of the clouds in the stratosphere, which is not already known but unique and unprecedented. We will add some comparison with the gound based observations of the high-altitude cloud diurnal cycle, even though theses observations suffer from little spatial coverage and high lidar attenuation by convective clouds in troposphere (see Authors' reponse to Referee #1).

- ✔ Referee #2:

Also, more references are needed. For example, it would be nice to have some reference of in-situ measurements (line 23 to 24) and of decreasing low-stratosphericumidity (line 41-43).

Authors' response:
Following Refere #2 recommendations, we added the following references:
(we got these recommendation during the initial round of review, so the changes are already present in the version of the manuscript posted for interactive discussion)
- for in-situ measurements, l.24: Thomas et al., 2002; Jensen et al., 2013; Frey et al., 2014
- for decreasing low-stratosphere humidity, l.46: Jensen et al., 2013.

The d~~D~~iurnal cycle of the clouds extending above the tropical tropopause observed by spaceborne lidar

Dauhut, Thibaut (1), Vincent Noel (2) and Iris-Amata Dion (3)

1. Max Planck Institute for Meteorology, Hamburg, Germany
2. Laboratoire d'Aérologie, Université de Toulouse, CNRS, UPS, Toulouse, France
3. CRNM, Météo-France – CNRS, Toulouse, 31057, France

Proposed for publication in Atmospheric Chemistry and Physics

Contact author : T. Dauhut thibaut.dauhut@mpimet.mpg.de

Abstract

The presence of clouds above the tropopause over tropical convection centers has so far been documented by spaceborne instruments that are either sun-synchronous, or insensitive to thin cloud layers. Here we document, for the first time through direct observation by spaceborne lidar, how the tropical cloud fraction evolves above the tropopause throughout the day. After confirming previous studies that found such clouds are most frequent above convection centers, we show that stratospheric clouds and their vertical extent above the tropopause follow a diurnal rhythm linked to convective activity. The diurnal cycle of the stratospheric clouds displays two maxima: one in the early night (19-20LT) and a later one (00-01LT). Stratospheric clouds extend up to 0.5-1km above the tropopause during nighttime, when they are the most frequent. The frequency and the vertical extent of stratospheric clouds is very limited during daytime, and when present they are found very close to the tropopause. Results are similar over the major convection centers (Africa, South America, Warm pool), with more clouds above land in DJF and less above ocean and JJA.

1. Scientific context and objectives

Low-stratospheric clouds impact the atmospheric system in several ways. First, their larger heating rate than the clear sky (Corti et al, 2006) increases the upward mass flux and fosters the large-scale upward transport of water above the tropopause. At the hour timescale, the cloud particles penetrating the stratosphere via overshooting convection leads, on the one hand, to a direct stratospheric humidification (Schoeberl et al., 2018; Dauhut et al., 2018). On the other hand, these particles can serve as support for ice-scavenging: under saturated conditions, the water vapor deposits on the particles, which grow and fall out (Corti et al., 2008), decreasing low-stratosphere humidity (Jensen et al., 2013). By all these effects the stratospheric clouds modulate the stratospheric water vapor concentrations (Iwasaki et al., 2015) and affect the overall dynamical structure near the tropopause (Corti et al., 2006), at timescales down to one hour. This is why it is important to understand the formation and the sub-daily evolution of such clouds.

The presence of ice clouds near the tropical tropopause has long been documented by in-situ measurements (e.g. Thomas et al., 2002; Jensen et al., 2013; Frey et al., 2014). Detecting occurrences of clouds extending above the tropopause by remote sensing requires documenting the vertical cloud profile with a fine resolution and a high sensitivity to optically thin clouds, which few instruments can reach. Lidar measurements are able to document such occurrences (e.g. Nee et al., 1998; Dupont et al., 2010; Gouveia et al., 2017), but for a long time were limited to local case studies. Dessler (2009) was the first to use the cloud detections by the CALIPSO lidar (Cloud-Aerosol Lidar Infrared Pathfinder Satellite Observations) to investigate how clouds extend above the tropopause on a global scale. Pan and Munchak (2011) refined the results by using an advanced tropopause dataset. Both studies found that clouds extending into the stratosphere are frequent above seasonal deep convection centers and rarely elsewhere, especially in midlatitudes. Both studies deplored that the fixed overpass local time of the CALIPSO dataset is far from the late afternoon, when land convection is at its maximum. More recently, Wang et al. (2019)

documented the presence of laminar cirrus in 10 years of CALIPSO data, and reported a non-negligible cloud amount above the tropopause. Because of the sun-synchronous orbit of CALIPSO, none of these studies were able to document the diurnal cycle of the stratospheric clouds.

~~Low stratospheric clouds impact the atmospheric system in several ways. First, their larger heating rate than the clear sky (Corti et al, 2006) increases the upward mass flux and fosters the large-scale upward transport of water above the tropopause. At the hour timescale, the cloud particles penetrating the stratosphere via overshooting convection leads, on the one hand, to a direct stratospheric humidification (Schoeberl et al., 2018; Dauhut et al., 2018). On the other hand, these particles can serve as support for ice scavenging: under saturated conditions, the water vapor deposits on the particles, which grow and fall out (Corti et al., 2008), decreasing low-stratosphere humidity (Jensen et al., 2013). By all these effects the stratospheric clouds modulate the stratospheric water vapor concentrations (Iwasaki et al., 2015) and affect the overall dynamical structure near the tropopause (Corti et al., 2006), at timescales down to one hour. This is why it is important to understand the formation and the sub-daily evolution of such clouds.~~

The diurnal evolution of the high-altitude cirrus clouds have been documented over some specific sites using ground-based lidars (Sassen et al. 2003; Dupont et al., 2010; Gouveia et al. 2017). Gouveia et al. (2017) documented the evolution of the integrated cloud fraction (no vertical distribution) over Amazonia, Sassen et al. (2003) documented the diurnal evolution of the composition of cirrus clouds over Salt Lake City, and Dupont et al. (2010) did the same over four observatories in France and in the United-States. However, using ground-based lidar to document optically thin clouds extending above the tropopause is difficult for two reasons: 1) as the studies based on CALIPSO observations show, these clouds occur primarily in regions where operational ground-based sites are absent or very few (Pacific ocean, equatorial Africa, South America), and 2) these clouds are mainly associated with deep convection, which implies the presence of optically thick cloud systems in the troposphere beneath that will make in most cases impossible the successful probing of optically thin clouds near the tropopause due to the attenuation of lidar signal. This explains why the ground-based lidars do not document the diurnal cycle of the stratospheric clouds with a satisfying spatial and temporal coverage.

Describing the diurnal evolution of the high-altitude clouds from a global perspective becomes possible with the CATS (Cloud-Aerosol Transport System) lidar operated from the International Space Station (ISS) between February 2015 and November 2017 (McGill et al., 2015). Thanks to the ISS non-synchronous orbit, CATS was able to probe the vertical cloud distribution of a particular region at different times of the day (not only at 0130 and 1330 Local Time like the instruments on CALIPSO). Aggregating CATS detections over a region of interest and over enough time provides a statistical overview of the diurnal evolution of cloud vertical profiles over that region (Noel et al., 2018). Our work aims at using CATS observations to describe and understand better the diurnal evolution of the cloud fraction in the tropical stratosphere.

Finding the processes responsible for the formation of tropical stratospheric clouds proves difficult, just like with high-tropospheric clouds (Reverdy et al., 2012). Two processes have been mainly proposed. Overshooting convection can lead to the injection of ice crystals into the stratosphere (Dauhut et al., 2018; Lee et al., 2018). Stratospheric cooling triggered by gravity waves (Pfister et al., 2010) could also lead to so-called cloud “in-situ” formation (Pan and Munchak, 2011). The ratio of stratospheric clouds that are formed in-situ has not been estimated yet. The current study does not provide further estimate, but by describing the spatio-temporal evolution of the stratospheric clouds, it highlights how important the convective activity is to drive the stratospheric cloudiness, and how the twice-daily sampling by lidars onboard sun-synchronous platforms can miss the highest and largest stratospheric cloud fraction over certain regions.

In this paper, we document for the first time the diurnal cycle of clouds above the tropopause in the Tropics, and the extent of their penetration in the stratosphere, thanks to the high vertical and temporal resolution of the cloud detection by the CATS~~ (Cloud Aerosol Transport System) spaceborne~~ lidar~~ (McGill et al., 2015)~~. After describing CATS cloud data, and the method to retrieve the tropopause heights used to detect clouds extending in the stratosphere (Sect. 2), we present maps of stratospheric clouds and document their diurnal cycle in regions of interest (Sect. 3). We then summarise our results and conclude (Sect. 4).

2. Data and Methods

2.1 CATS Cloud data

Between February 2015 and November 2017, ~~T~~the CATS lidar ~~operated from the International Space Station (ISS) between February 2015 and November 2017. It~~ reported profiles at a vertical resolution of 60m every 350m along-track, with an average repeat cycle of nearly 3 days (Yorks et al., 2016). ~~Thanks to the ISS non synchronous orbit, CATS was able to probe the vertical cloud distribution of a particular region at different times of the day. Aggregating CATS detections over a region of interest and over enough time provides a statistical overview of the diurnal evolution of cloud vertical profiles over that region (Noel et al., 2018).~~ CATS Level 2 Operational layer files (L2O files, Palm et al., 2016) describe altitudes where cloud layers were detected within profiles of backscatter coefficients measured at 1064nm by the CATS lidar (Pauly et al., 2019), averaged 5km along-track. We considered all such files over the CATS operation period ~~(February 2015 to November 2017)~~ and inspected each 5-km profile within. For profiles located in the Tropics (30S-30N), we inspected each atmospheric layer therein identified as a cloud layer according to the CATS layer type information. As in Noel et al. (2018), we considered layers with a Feature Type Score above 6, to avoid any possibly mislabeled aerosol layers. We flagged the cloud layers with a top altitude above the tropopause. Since any CATS L2O layer entirely above the tropopause is labelled as an aerosol layer (like in CALIPSO, Pan and Munchak, 2011), our study will not include clouds with their base in the stratosphere.

Davis et al (2010) noted that lidars in space may miss the thinnest subvisible cirrus clouds, but with enough spatial averaging optical depths near 0.001 can be detected (Martins et al., 2011). Lidar cloud detections also suffer from a lower sensitivity in the presence of sunlight, which induces significant additional noise in the lidar signal, but climatologies are still relevant (Noel et al., 2018).

2.2. Tropopause Heights

To obtain the tropopause height, we considered profiles of temperature and pressure from the ERA-5 reanalysis dataset (Albergel et al., 2018). These profiles are available every 6 hours, on 37 vertical levels and a 0.25° x 0.25° horizontal grid. Such profiles in ERA-5 reanalysis agree well with observations in the high tropical troposphere (Podglajen et al. 2014). Using these profiles, we computed the vertical lapse rate profile (as in Reichler et al. 2003), and interpolated it on a 100-m vertical grid. We then applied the WMO criteria defining the presence of a tropopause -- i.e. the lowest altitude at which the lapse rate falls below 2°C/km, provided the lapse-rate between this ~~level~~altitude and all higher ~~levels~~altitudes within 2 km does not exceed 2°C/km (WMO, 1957). Following the WMO definition, we also allowed for the possibility of a second tropopause if the lapse rate exceeds 3°C/km at least 1 km above the first tropopause. In such a case, we started to look for another tropopause above. To limit computation overhead we constrained the search below 22 km. Using the WMO tropopause definition further allows us to compare our results to previous efforts based on CALIPSO database that used the same definition (Pan and Munchak, 2011).

2.3 Stratospheric cloud detection

For a given CATS 5-km profile (Sect. 2.1), we identified the ERA-5 tropopause height (Sect. 2.2) closest in time and location. Given the 6-hour time resolution of the ERA-5 reanalysis, there is at most 3 hours difference between the observation time and the thermodynamic information used to retrieve the tropopause height. We used the cloud information contained in the 5-km profiles in two ways. First, in 2°x5° lat-lon bins we counted how many profiles contained a cloud extending above the tropopause, compared to the total number of profiles in the bins. Aggregating such numbers observed in JJA and DJF over the CATS operation period produced seasonal maps of above-tropopause cloud amounts (Sect. 3.1). Second, from each CATS 5-km profile we built a vertical cloud mask, using the tropopause height as the vertical reference and considering clouds that extend above it. Within regions chosen based on the seasonal maps, we aggregated such cloud masks over the same periods as above, keeping also track of the local time of observation for the considered mask. This produced regional vertical cloud fraction profiles above the tropopause, one profile for each local time of observation (Sect. 3.2).

3. Results

3.1 Stratospheric cloud distributions

Figure 1 shows the fraction of CATS profiles in which a cloud is detected above the tropopause, in all DJF (top) and JJA (bottom) months of CATS operation.

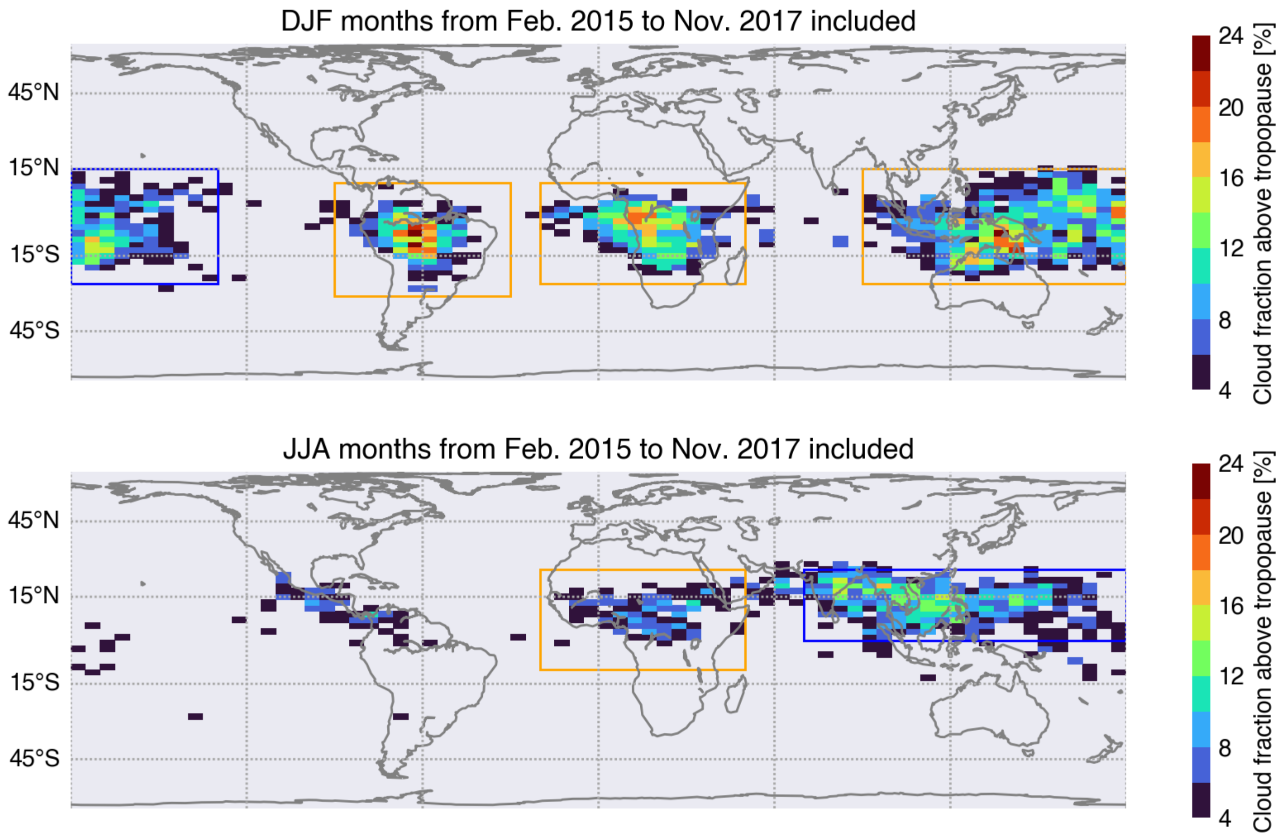

Figure 1*: Tropical low-stratosphere cloud fraction for (top) DJF and (bottom) JJA CATS measurements between Feb 2015 and Nov 2017, calculated by considering all profiles in 2°x5° lat-lon boxes. The rectangles are the regions in which cloud detections are aggregated in the rest of the study. In DJF, from left to right : West Pacific (25S-15N, 180W-130W), South America (30S-10N, 90W-30W), Equatorial Africa (25S-10N, 20W-50E), and South Warm Pool (25S-15N, 90E-180E). In JJA, from left to right : Central Africa (10S-25N, 20W-50E), North Warm Pool (0-25N, 70E-180E). Only detections in the ±30° region are shown here. In the rest of the study, we considered profiles over ocean in blue boxes and profiles over land in orange boxes.*

Figure 1 shows that clouds in the tropical stratosphere are mostly detected over continents (South America, Equatorial Africa and land masses in South Warm Pool in DJF; Central America, Central Africa and land masses in North Warm Pool in JJA). The cloud fraction in the lower stratosphere is largest in DJF, up to 24% over central Amazonia and coastal areas in South Warm Pool, and up to 20% over Equatorial Africa. It is significantly lower in JJA, up to 12% over Africa and 16% over the North Warm Pool, even though the lowermost stratosphere (380-420 K potential temperature) is moister in JJA than in DJF (cf. e.g. Fig. 8c in Fueglistaler et al, 2009). Several factors may contribute to this seasonal variation: the density and strength of the convective systems (Liu and Zipser, 2005), their propensity to propagate or to be stationary (Houze et al, 2015), the activity and efficiency of the in situ formation processes (Jensen et al., 2001; Jensen and Pfister, 2004).

The spatio-temporal distribution of the stratospheric clouds is in very good agreement with the 4-year climatology of Pan and Munchak (2011) from CALIPSO observations. The DJF distribution also matches very well the ~~CALIOP~~CALIPSO cirrus detection at 100 hPa reported by Wang et al (2019) for January 2009. We report though lower cloud frequencies than Wang et al. (2019) which can be explained that we investigate slightly higher ~~levels~~altitudes. Both CATS and CALIPSO datasets find 1) significantly weaker stratospheric cloud fraction in JJA than in DJF, and 2) near-zero stratospheric clouds in the subtropics. These results are also consistent with the CALIPSO cloud fractions near 16km reported by Schoeberl et al. (2019). Since those studies consider cloud detections derived from a spaceborne lidar instrument, over several years for most, their good agreement suggests that the CATS stratospheric cloud detections at 1064 nm are as reliable as the CALIPSO ones at 532 nm. A first conclusion of our results is therefore that CATS measurements strongly support the findings of all other studies using detections of high clouds from CALIPSO data.

Our CATS results are also in very good agreement with the distributions of clouds near the tropopause from other space instruments: 2006-2007 HIRDLS (High Resolution Dynamics Limb Sounder) reported by Massie et al. (2010), 2006-2014 CloudSat observations (Kim et al., 2018), and the pioneering 1989 passive Stratospheric Aerosol and Gas Experiment (SAGE) II observations (Jensen et al., 1996). Besides the specificity in the cloud detection method employed by each instrument (occultation for HIRDLS and SAGE II, radar backscattering for CloudSat), the little differences between the distributions mostly come from the year-to-year variability. Larger differences can be found with the distributions of clouds penetrating the tropical tropopause derived from the 1998-2000 and 2002-2003 observations by the Tropical Rainfall Measuring Mission (TRMM) Precipitation Radar (Liu and Zipser, 2005). The densities of overshooting systems with tops in the lower stratosphere (on which Liu and Zipser (2005) focused rather than all stratospheric clouds) are remarkably larger in Central America and Central Africa than over the Warm Pool. Since TRMM precipitation radar reflectivities are less sensitive to thin ice particles than CATS and CALISPO lidars, we can interpret this difference by the fact that the American and African systems, though frequently overshooting the stratosphere, produce less thin stratospheric clouds than the Asian systems, or other in-situ processes (like gavity wave cooling) are more efficient to produce stratospheric clouds over Asia than America and Africa.

Comparing our CATS results to the distributions of clouds at 90 hPa/17 km retrieved from HIRDLS and CALIPSO for 2006-2007 by Massie et al. (2010), we also find good agreement in JJA but larger differences in DJF. In CALIPSO and HIRDLS, the maxima are over the West to Central Pacific and the convective spot in South America is shifted West towards the Pacific. This difference can be explained by the annual variability: in DJF 2006-2007 the Southern Oscillation Index indicates rather El Nino conditions, like in DJF 2015-2016 but in contrast with DJF 2016-2017, both being included in our study.

Our results match the overall distributions of cloud tops higher than 17 km retrieved from CloudSat 8-year observations by Kim et al. (2018). In particular, the Warm Pool exhibits the largest area with significant stratospheric cloud fraction, both during DJF and JJA. Little differences between CATS and CloudSat data sets appear in DJF: in the CATS dataset the South America show slightly larger cloud fraction than the Equatorial Africa; in the CloudSat dataset, the largest cloud fractions over Africa are located more south-east (Great Lakes and Madagascar straight). Note that the CloudSat radar samples convection at 1:30 am and pm, potentially missing some continental convective systems. These differences might also be due to the different periods considered: 2006-2014 for CloudSat versus 2015-2017 for CATS. On the contrary, our results contrast with Liu and Zipser (2005) distributions derived from the TRMM Precipitation Radar, where the densities of overshooting systems with tops in the lower stratosphere are remarkably larger in Central America and Central Africa than over the Warm Pool. Since TRMM precipitation radar reflectivities are less sensitive to thin ice particles than CATS and CALIOP lidars, we can interpret this difference by the fact that the American and African systems, though frequently overshooting the stratosphere, produce less stratospheric clouds than the Asian systems.

Finally, our results agree well with the pioneering work of Jensen et al. (1996) who used passive SAGE II observations at 17.5 km for 1989: their cloud fractions are larger because the considered level is closer to the cold point tropopause but the geographical distributions are very close to ours. The differences are: in DJF they observe more clouds over the Atlantic but less over South America, in JJA they observe less clouds over the West Pacific. These differences may be due to the year-to-year variability. The SAGE II instrument relies on a solar occultation method, completely different from the active lidar observation by CATS and CALIPSO.

3.2 Diurnal cycle of cloud fractions in the tropical stratosphere

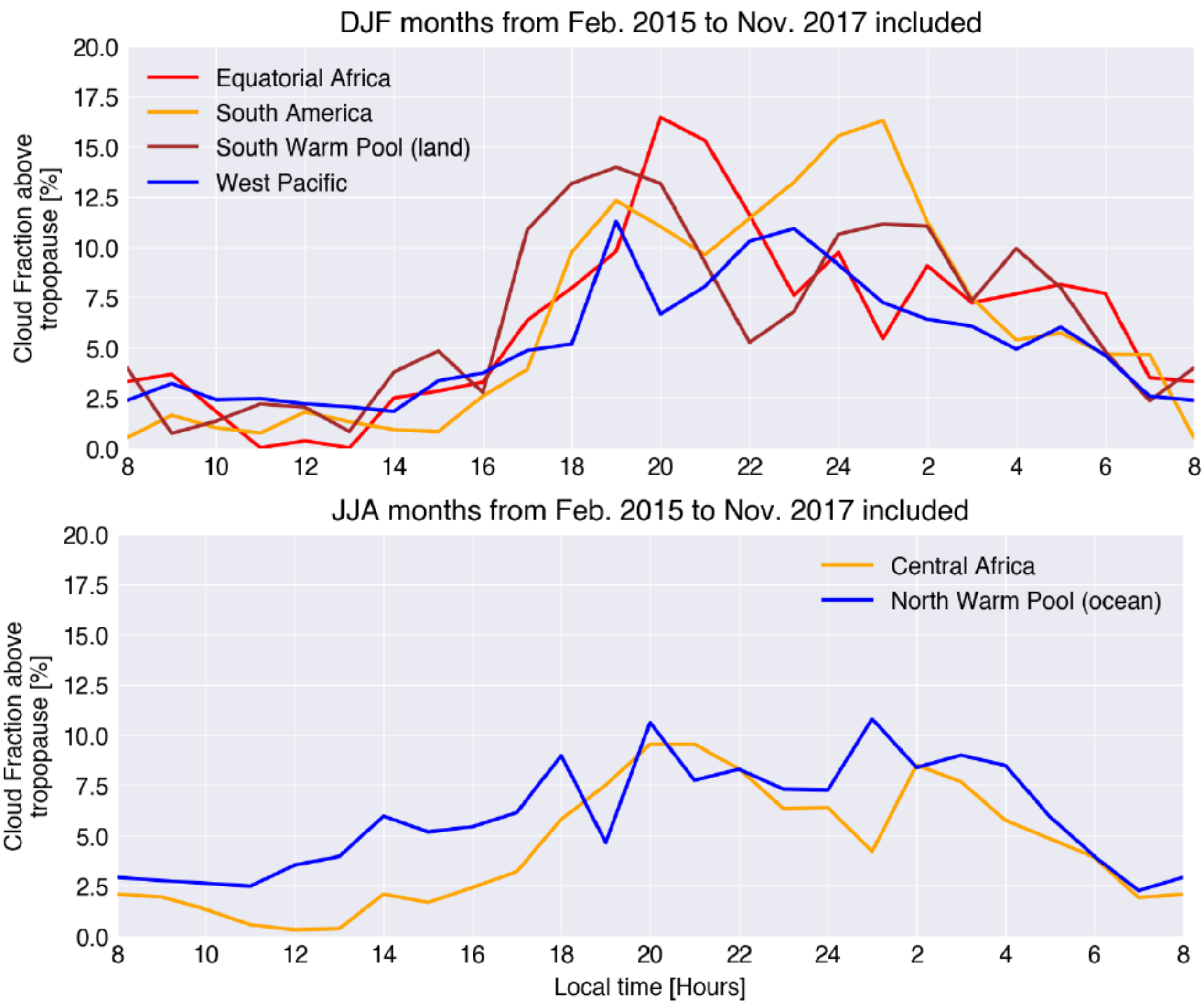

***Figure 2**: Diurnal cycle of stratospheric cloud fraction, by tropical region as in Fig. 1, averaged over DJF (top) and JJA (bottom).*

In contrast with the previous studies, the CATS dataset allows us to analyse the diurnal cycle of the cloud fraction in stratosphere. The cloud fraction at regional scale shows a consistent diurnal cycle, robust over the different regions identified in the previous section (Fig. 2). In particular and in contrast to the diurnal cycle of surface precipitation, there is no land-ocean difference. All exhibit a pronounced minimum about 2-4 % during the day time, from 7 to 16 LT. ~~They all present a first maximum at 19 or 20 LT (early-night peak), up to 18% over South Warm Pool; over North Warm Pool this peak is slightly later. For all regions except South America, this maximum is the largest cloud fraction of the day. All regions also present a secondary peak at 0 or 1 LT (2 LT for Central Africa), up to 16% over South America.~~They all present a first maximum at 19 or 20 LT (early-night peak), up to 16.5% over Equatorial Africa. For all regions except South America and North Warm Pool, this maximum is the largest cloud fraction of the day. All regions also present a second peak (late-night peak) at 0 or 1 LT (23 LT for West Pacific and 2 LT for Central Africa), up to 16.5% over South America. The midnight peak over Equatorial Africa is less clear than over the other regions because of the large variations between 23 and 3 LT. The capability of a longer dataset to produce a clearer signal remains to be investigated.

The cirrus clouds observed over Amazonia by gound-based lidar (Gouveia et al., 2017) shows a very similar diurnal cycle: a first peak in the early night (at 18-19 LT), a second peak later in the night (at 2-3 LT). Although Gouveia et al. (2017) do not consider the cloud above the tropopause only, their distinction between subvisible, thin and opaque cirrus indicates that the opaque cirrus are predominant during the early night (18-21 LT) and the thin cirrus (and subvisible ones during the dry season) dominate during the later night (from 0 and 2 LT onward, in wet and dry seasons, respectively).

The very deep convection transports cloudy air masses beyond the tropopause via overshoots and then directly contributes to the stratospheric cloud fraction (Dauhut et al., 2016 and 2018). The diurnal cycle of the stratospheric cloud fraction observed by CATS can at the first order be explained by the diurnal cycle of very deep convection over land (Liu and Zipser, 2005), especially (i) the minimal value during daytime, and (ii) the first peak in the early evening. This first peak occurs with a delay of 3 to 4 hours compared to the very deep convection maximum. As the dataset used by Liu and Zipser (2005) is more sensitive to overshoots freshly developed into the stratosphere, this delay can be explain by the subsequent horizontal expansion of the overshoots and their spread by the winds (Dauhut et al., 2018; Lee et al., 2019). The convective generation of gravity waves, that produce transient cooling off the convective centres and in some conditions trigger cloud formation, can also contribute to the increase of the stratospheric cloud fraction after the maximum of the very deep convection, and then explain the delay of the first peak and potentially the second peak. It may also explain the similar diurnal cycle over the ocean regions, either close (South Warm Pool Ocean) or remote (West Pacific) from land masses. This process remains to be investigated.

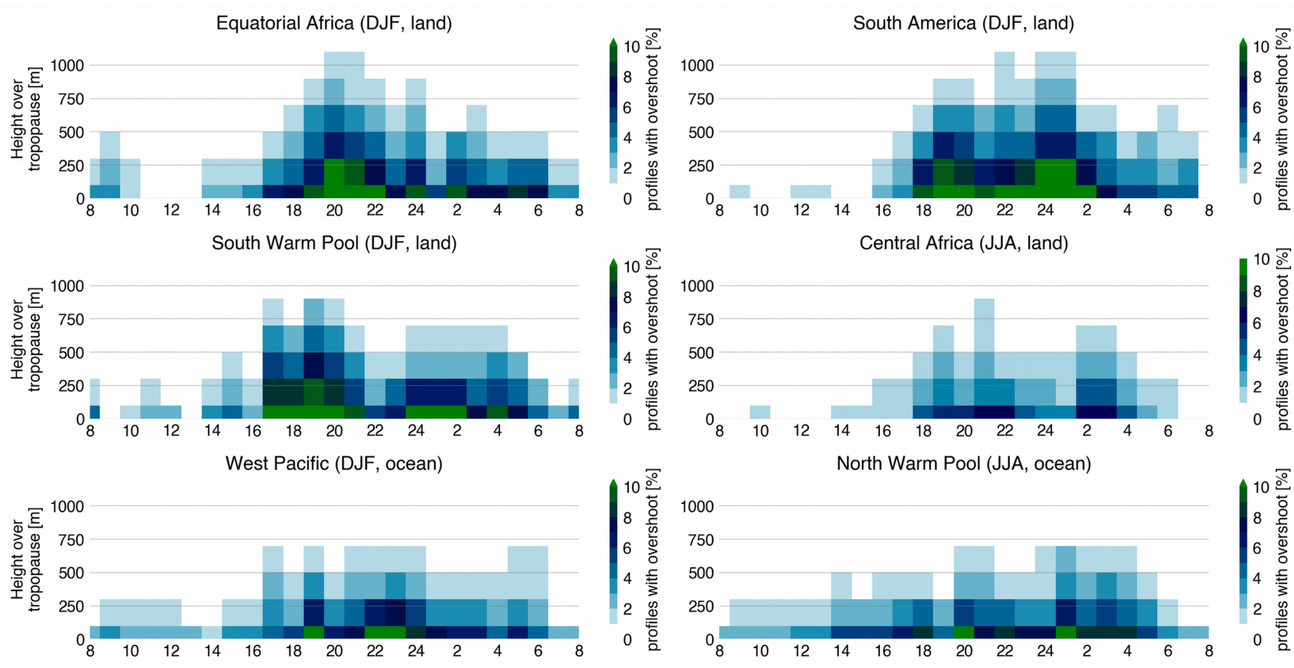

__Figure 3:__ Diurnal cycle of cloud fraction as a function of height above the tropopause, by tropical region as in Fig. 1, in DJF (top 2 rows) and JJA (bottom row).

Figure 3 shows how far above the tropopause the clouds extend, depending on the local time in each tropical region (Sect. 3.1). Some regions are considered in DJF, others in JJA, because the stratospheric cloud distribution changes throughout the year (Fig. 1), following the ITCZ position. Patterns appear very consistent in all the regions considered. In all regions the largest cloud fractions are found near the tropopause, with few clouds extending higher. Cloud fractions extend relatively high (up to 1km above the tropopause) during the early night. The first peak of cloud fraction, near 19-20LT (Fig. 2), is associated with the all-day maximum of cloud vertical extent, with clouds in 5% of profiles reaching 1km above the tropopause in DJF regions. During the rest of the night (after 00 LT) clouds are still present but extend less high (except over South America). During daytime (0600-1800) clouds appear very close to the tropopause. Cloud fractions are overall much smaller in JJA (max 5-10%, bottom row) than in DJF (max 10-12%, rows 1 and 2).

In addition to describing the evolution of the stratospheric cloud cover at hourly timescales, these observations help interpret observations with limited temporal sampling (Noel et al., 2018). The Microwave Limb Sounder (MLS), like CALIPSO and all other instruments onboard platforms of the A-Train, samples the atmosphere at 01:30 and 13:30 LT, providing one single night and one single day observation. Some authors (e.g., Dion et al., 2019) attempt to retrieve the diurnal cycle of the observed water contents in the tropopause region, combining MLS observations with higher temporal resolution observation of convective activity based on TRMM observation of precipitation. Dion et al. (2019) assumed an in-phase relationship between precipitation and ice water content in the upper troposphere and at the tropopause level. For the stratospheric ice water content, MLS data still provides a too low signal-to-noise ratio. For future investigations, our results indicate that the stratospheric cloud fraction at 13:30 LT is, whatever the region, close to the minimal value of its diurnal cycle, whereas at 01:30 LT it is more typical of the second~~ary~~ maximum. Carminati et al. (2014) investigated, from MLS measurements between 2005 and 2012, the differences between day and night ice water contents in the upper troposphere and the tropopause level. Unlike the stratospheric cloud fraction, tropopause ice water contents are larger at 13:30 LT than at 01:30 LT over Equatorial Africa during DJF, Central Africa during JJA, and over South America during both seasons. A possible explanation to reconcile our results is that tropopause ice water content is more sensitive to fresh convective activity (very deep convection occurrence) whereas the stratospheric cloud cover is more sensitive to the diffusion of the injected ice in the stratosphere.

4. Conclusion

Our results show how clouds in the tropical stratosphere are strongly concentrated above deep convection centers, are almost absent in subtropical regions, are more frequent in DJF than JJA, and over land than over ocean. In addition to these results, which are consistent with most previous studies, we also show that both the cloud fraction and its extension above the tropopause follow a diurnal rhythm with a maximum during the early nighttime and a near-zero minimum during daytime. During daytime, the stratospheric clouds are limited to the first hundred meters above the tropopause. During nighttime, significant average cloud fraction is found up to 1 km above the tropopause. A second~~ary~~ maximum of stratospheric cloud fraction is observed over all regions, generally little after midnight. These results highlight how much the evolution of stratospheric clouds can be undersampled by other spatial instruments restricted to 01:30 and 13:30 LT, that then miss for instance the first maximum and the deepest development of stratospheric clouds in the early night. ~~Further investigation is necessary to identify the processes driving~~ The very deep convective activity over tropical lands drives most of this diurnal cycle, and leads~~ing~~ in particular to the minimal stratospheric cloud fraction during daytime and the second~~ary~~ peak during nighttime, both consistent over all regions. Further investigation is though necessary to describe how convection contributes to this diurnal cycle, and to assess the role of other processes leading to stratospheric cloud formation like the gravity waves. Finally further research is needed to understand why the timing of this diurnal cycle is very similar over land and over ocean.

Author contribution. TD and VN designed the data analyses and VN carried them out. TD prepared the manuscript with contributions from VN and ID.

Acknowledgments. The authors would like to thank B. Legras (LMD) and F. Pantillon (LA) for useful discussions on the quality of tropopause altitudes from various sources. They also thank

NASA EarthData for access to CATS data and ECMWF for access to the ERA5 reanalysis dataset. This research was supported by the IDEX TEASAO project. Primary data and scripts used in the analysis and other supplementary information that may be useful in reproducing the author's work are archived by the Max Planck Institute for Meteorology and can be obtained by contacting publications@mpimet.mpg.de

References

- Albergel, C., Dutra, E., Munier, S., Calvet, J.-C., Munoz-Sabater, J., de Rosnay, P., and Balsamo, G.: ERA-5 and ERA-Interim driven ISBA land surface model simulations: which one performs better?, Hydrol. Earth Syst. Sci., 22, 3515-3532, https://doi.org/10.5194/hess-22-3515-2018, 2018.
- Carminati, F., Ricaud, P., Pommereau, J.-P., Rivière, E., Khaykin, S., Attié, J.-L., and Warner, J.: Impact of tropical land convection on the water vapour budget in the tropical tropopause layer, Atmos. Chem. Phys., 14, 6195-6211, https://doi.org/10.5194/acp-14-6195-2014, 2014.
- Corti, T., Luo, B. P., Fu, Q., Vömel, H., and Peter, T.: The impact of cirrus clouds on tropical troposphere-to-stratosphere transport, Atmos. Chem. Phys., 6, 2539-2547, https://doi.org/10.5194/acp-6-2539-2006, 2006.
- Corti, T., et al.: Unprecedented evidence for deep convection hydrating the tropical stratosphere, Geophys. Res. Lett., 35, L10810, https://doi.org/10.1029/2008GL033641, 2008.
- Dauhut, T., Chaboureau, J., Haynes, P.H., and Lane, T.P.: The Mechanisms Leading to a Stratospheric Hydration by Overshooting Convection, J. Atmos. Sci., 75, 4383–4398, https://doi.org/10.1175/JAS-D-18-0176.1, 2018.
- Davis, S., et al.: In situ and lidar observations of tropopause subvisible cirrus clouds during TC4, J. Geophys. Res., 115, D00J17, https://doi.org/doi:10.1029/2009JD013093, 2010.

- Dessler, A. E.: Clouds and Water Vapor in the Northern Hemisphere Summertime Stratosphere, 114, J. Geophys. Res., 114, https://doi.org/10.1029/2009JD012075, 2009.
- Dion, I. A., Ricaud, P., Haynes, P., Carminati, F., and Dauhut, T.: Ice injected into the tropopause by deep convection – Part 1: In the austral convective tropics, Atmos. Chem. Phys., 19(9), 6459-6479, https://doi.org/10.5194/acp-19-6459-2019, 2019.
- Frey, W., Borrmann, S., Fierli, F., Weigel, R., Mitev, V., Matthey, R., Ravegnani, F., Sitnikov, N. M., Ulanovsky, A., and Cairo, F.: Tropical deep convective life cycle: Cb-anvil cloud microphysics from high-altitude aircraft observations, Atmos. Chem. Phys., 14, 13223-13240, https://doi.org/10.5194/acp-14-13223-2014, 2014.
- Houze, R. A., Rasmussen, K. L., Zuluaga, M. D., and Brodzik, S. R.: The variable nature of convection in the tropics and subtropics: A legacy of 16 years of the Tropical Rainfall Measuring Mission satellite, Rev. Geophys., 53, 994–1021, https://doi.org/10.1002/2015RG000488, 2015.
- Iwasaki, S., Luo, Z. J., Kubota, H., Shibata, T., Okamoto,H., and Ishimoto, H.: Characteristics of cirrus clouds in the tropical lower stratosphere, Atmos. Res., 164–165, 358–368, https://doi.org/10.1016/j.atmosres.2015.06.009, 2015.
- Jensen, E. J., Toon, O. B., Selkirk, H. B., Spinhirne, J. D., and Schoeberl, M. R.: On the formation and persistence of subvisible cirrus clouds near the tropical tropopause, J. Geophys. Res., 101(D16), 21361– 21375, https://doi.org/10.1029/95JD03575, 1996.
- Jensen, E. J., Pfister, L., Ackerman, A. S., Tabazadeh, A., and Toon, O. B.: A conceptual model of the dehydration of air due to freeze-drying by optically thin, laminar cirrus rising slowly across the tropical tropopause, J. Geophys. Res., 106(D15), 17237– 17252, https://doi.org/10.1029/2000JD900649, 2001.
- Jensen, E., and Pfister, L.: Transport and freeze-drying in the tropical tropopause layer, J. Geophys. Res., 109, D02207, https://doi.org/10.1029/2003JD004022, 2004.
- Jensen, E. J., Diskin, G., Lawson, R. P., Lance, S., Bui, T. P., Hlavka, D., McGill, M., Pfister, L., Toon, O. B., and Gao, R.: Ice nucleation and dehydration in the Tropical Tropopause Layer, PNAS, 110, 2041–2046, doi:10.1073/pnas.1217104110, 2013.

- Kim, J., Randel, W. J., and Birner, T.: Convectively driven tropopause-level cooling and its influences on stratospheric moisture, J. Geophys. Res.-Atmos., 123, 590– 606, https://doi.org/10.1002/2017JD027080, 2018.
- Lee, K.-O., Dauhut, T., Chaboureau, J.-P., Khaykin, S., Krämer, M., and Rolf, C.: Convective hydration in the tropical tropopause layer during the StratoClim aircraft campaign: Pathway of an observed hydration patch, Atmos. Chem. Phys.~~ Discuss.~~, https://doi.org/10.5194/acp-2018-1114, ~~preprint, 2018~~2019.
- Liu, C., and Zipser, E. J.: The global distribution of largest, deepest, and most intense precipitation systems, Geophys. Res. Lett., 42, 3591– 3595, https://doi.org/10.1002/2015GL063776, 2005.
- Martins, E., Noel, V., and Chepfer, H.: Properties of cirrus and subvisible cirrus from nighttime Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP), related to atmospheric dynamics and water vapor, J. Geophys. Res., 116, D02208, https://doi.org/10.1029/2010JD014519, 2011.
- Massie, S. T., Gille, J., Craig, C., Khosravi, R., Barnett, J., Read, W., and Winker, D.: HIRDLS and CALIPSO observations of tropical cirrus, J. Geophys. Res., 115, D00H11, https://doi.org/10.1029/2009JD012100, 2010.
- McGill, M. J., Yorks, J. E., Scott, V. S., Kupchock, A. W., and Selmer, P. A.: The Cloud-Aerosol Transport System (CATS): A technology demonstration on the International Space Station, Proc. SPIE 9612, Lidar Remote Sensing for Environmental Monitoring XV, 96120A, https://doi.org/10.1117/12.2190841, 2015.
- Nee, J. B., Len, C. N., Chen, W. N., and Lin, C. I.: Lidar observation of the cirrus cloud in the tropopause at Chung Li (25°N, 121°E), J. Atmos. Sci., 55, 2249– 2257, 1998.
- Noel, V., Chepfer, H., Chiriaco, M., and Yorks, J.: The diurnal cycle of cloud profiles over land and ocean between 51° S and 51° N, seen by the CATS spaceborne lidar from the International Space Station, Atmos. Chem. Phys., 18, 9457-9473, https://doi.org/10.5194/acp-18-9457-2018, 2018.

- Palm, S. P., Hlavka, D. L., Selmer, P., and Pauly, R.: the Cloud Aerosol Transport System (CATS) Data Product Catalog release 3.0, available at: https://cats.gsfc.nasa.gov/media/docs/CATS_Data_Products_Catalog.pdf (last access: 23 January 2018), 2016.
- Pan, L. L., and Munchak, L. A.: Relationship of Cloud Top to the Tropopause and Jet Structure from CALIPSO Data, J. Geophys. Res., 116, D12201, https://doi.org/10.1029/2010JD015462, 2011.
- Pauly, R. M., Yorks, J. E., Hlavka, D. L., McGill, M. J., Amiridis, V., Palm, S. P., Rodier, S. D., Vaughan, M. A., Selmer, P. A., Kupchock, A. W., Baars, H., and Gialitaki, A.: Cloud Aerosol Transport System (CATS) 1064 nm Calibration and Validation, Atmos. Meas. Tech. Discuss., https://doi.org/10.5194/amt-2019-172, 2019
- Pfister, L., Selkirk, H. B., Starr, D. O., Rosenlof, K., and Newman, P.A.: A meteorological overview of the TC4 mission, J. Geophys. Res., 115, D00J12, https://doi.org/10.1029/2009JD013316, 2010.
- Podglajen, A., Hertzog, A., Plougonven, R., and Žagar, N.: Assessment of the accuracy of (re)analyses in the equatorial lower stratosphere, J. Geophys. Res. Atmos., 119, 11166–11188, https://doi.org/10.1002/2014JD021849, 2014.
- Reichler, T., Dameris, M., and Sausen, R.: Determining the Tropopause Height from Gridded Data, Geophys. Res. Lett., 30, 2042, https://doi.org/10.1029/2003GL018240, 2003.
- Reverdy, M., Noel, V., Chepfer, H., and Legras, B.: On the origin of subvisible cirrus clouds in the tropical upper troposphere, Atmos. Chem. Phys., 12, 12081-12101, https://doi.org/10.5194/acp-12-12081-2012, 2012.
- Schoeberl, M. R., Jensen, E. J., Pfister, L., Ueyama, R., Wang, T., Selkirk, H., et al.: Water vapor, clouds, and saturation in the tropical tropopause layer. *J. Geophys. Res.-Atmos.*, 124, 3984– 4003, https://doi.org/10.1029/2018JD029849, 2019.
- Thomas, A., Borrmann, S., Kiemle, C., Cairo, F., Volk, M., Beuermann, J., Lepuchov, B., Santacesaria, V., Matthey, R., Rudakov, V., Yushkov, V., MacKenzie, A. R., and Stefanutti,

L.: In situ measurements of background aerosol and subvisible cirrus in the tropical tropopause region, J. Geophys. Res., 107, 4763, doi:10.1029/2001JD001385, 2002.

- Wang, T., Wu, D. L., Gong, J., and Tsai, V.: Tropopause laminar cirrus and its role in the lower stratosphere total water budget. *J. Geophys. Res.-Atmos.*, 124, 7034– 7052, https://doi.org/10.1029/2018JD029845, 2019.
- Yorks, J. E., McGill, M. J., Palm, S. P., Hlavka, D. L., Selmer, P. A., Nowottnick, E. P., Vaughan, M. A., Rodier, S. D., and Hart, W. D.: An overview of the CATS level 1 processing algorithms and data products, Geophys. Res. Lett., 43, 4632– 4639, https://doi.org/10.1002/2016GL068006, 2016.