

1 **Heterogeneous Formation of Particulate Nitrate under Ammonium-**
2 **rich Regime during the High PM_{2.5} Events in Nanjing, China**

3 **Yu-Chi Lin^{1,2,3}, Yan-Lin Zhang^{1,2,3*}, Mei-Yi Fan^{1,2,3}, Mengying Bao^{1,2,3}**

4 ¹. *Yale-NUIST Center on Atmospheric Environment, International Joint Laboratory on*
5 *Climate and Environment Change, Nanjing University of Information Science and*
6 *Technology, Nanjing, 210044, China.*

7 ². *Key Laboratory Meteorological Disaster; Ministry of Education & Collaborative*
8 *Innovation Center on Forecast and Evaluation of Meteorological Disaster, Nanjing*
9 *University of Information Science and Technology, Nanjing, 210044, China.*

10 ³. *Jiangsu Provincial Key Laboratory of Agricultural Meteorology, College of Applied*
11 *Meteorology, Nanjing University of Information Science & Technology, Nanjing*
12 *210044, China.*

13
14 *Corresponded to Yan-Lin Zhang (dryanlinzhang@outlook.com;*
15 *zhangyanlin@nuist.edu.cn)*

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17 **ABSTRACT**

18 Particulate nitrate (NO₃⁻) not only influences regional climates but also contributes to
19 the acidification of terrestrial and aquatic ecosystems. In 2016 and 2017, four
20 intensive online measurements of water-soluble ions in PM_{2.5} were conducted in
21 Nanjing City to investigate the potential formation mechanisms of particulate nitrate.
22 During the sampling periods, NO₃⁻ was the most predominant species, accounting for
23 35 % of the total water-soluble inorganic ions, followed by SO₄²⁻ (33 %) and NH₄⁺
24 (24 %). Significant enhancements of nitrate aerosols in terms of both absolute
25 concentrations and relative abundances suggested that NO₃⁻ was a major contributing

26 species to high-PM_{2.5} events (hourly PM_{2.5} \geq 150 $\mu\text{g m}^{-3}$). High NO₃⁻
27 concentrations mainly occurred under NH₄⁺-rich conditions, implying that the
28 formation of nitrate aerosols in Nanjing involved NH₃. During the high-PM_{2.5} events,
29 the nitrogen conversion ratios (Fn) were positively correlated with the aerosol liquid
30 water content (ALWC, R > 0.72, p < 0.05). Meanwhile, increasing NO₃⁻
31 concentrations regularly coincided with increasing ALWC and decreasing Ox (Ox =
32 O₃ + NO₂). These results suggested that the heterogeneous reaction was probably a
33 major mechanism of nitrate formation during the high-PM_{2.5} events. Moreover, the
34 average production rate of NO₃⁻ by heterogeneous processes was estimated to be 12.6
35 % h⁻¹ (4.1 $\mu\text{g m}^{-3} \text{ h}^{-1}$), which was much higher than that (2.5 % h⁻¹, 0.8 $\mu\text{g m}^{-3} \text{ h}^{-1}$) of
36 gas-phase reactions. This can also explain the abrupt increase of nitrate concentrations
37 during the high PM_{2.5} events. Utilization of ISORROPIA II model, we found that
38 nitrate aerosol formation in Nanjing during the high-PM_{2.5} events was HNO₃-limited.
39 This indicated that the control of NO_x emissions will be able to efficiently reduce
40 airborne nitrate concentrations and improve the air quality in this industrial city.
41 Keywords: Nitrate aerosols, nitrogen conversion ratios, NH₄⁺-rich regime, Hydrolysis
42 of N₂O₅, Nitrate production rate

43

44 1. Introduction

45 Due to the rapid growth of industrialization and urbanization, particulate matter
46 (PM) pollution has become a severe problem in China in recent years (Chan and Yao,
47 2008; Zhang and Cao, 2015). Fine mode particles (PM_{2.5}, with aerodynamic diameters
48 less than 2.5 μm) exhibit smaller sizes and contain many toxins emitted from
49 anthropogenic emissions (Huang et al., 2018). PM_{2.5} easily penetrates the upper
50 respiratory tract and is deposited into the human body, causing serious threats to
51 human health. Numerous previous studies have proven that people exposed to high

52 PM_{2.5} concentrations show increased risks of respiratory illness, cardiovascular
53 diseases and asthma (Brauer et al., 2002; Defino et al., 2005), resulting in an increase
54 of mortality (Nel, 2005).

55 Secondary inorganic aerosols (SIA), including sulfate (SO₄²⁻), nitrate (NO₃⁻) and
56 ammonium (NH₄⁺), are major constituents of PM_{2.5}, accounting for 25 - 60 % of the
57 PM_{2.5} mass in urban cities of China (Huang et al., 2014a; Wang et al., 2018; Yang et
58 al., 2005; Ye et al., 2017; Zhao et al., 2013; Zhou et al., 2018). Among these species,
59 SO₄²⁻ and NO₃⁻ are acidic ions which tend to be neutralized by NH₄⁺. Previously,
60 many studies suggested that SO₄²⁻ dominated SIA in urban cities of China (Kong et
61 al., 2014; Tao et al., 2016; Yang et al., 2005; Yao et al., 2002; Zhao et al., 2013). In
62 recent years, the Chinese government reduced its anthropogenic emissions by 62 %
63 and 17 % for SO₂ and NO_x, respectively (Zheng et al., 2018). This revealed that the
64 reduction efficiency of SO₂ emissions were much higher than those of NO_x.
65 Consequently, nitrate has become the dominant species of SIA, especially during PM
66 haze events (Wang et al., 2018; Wen et al., 2015; Zou et al., 2018).

67 In the atmosphere, ammonium nitrate (NH₄NO₃) is a major form of nitrate
68 aerosols in fine mode particles. NH₄NO₃ is a semi-volatile species which partitions
69 from the particle phase into the gas phase under high-temperature (T) conditions. It
70 deliquesces when the ambient relative humidity (RH) is higher than its deliquescence
71 relative humidity (DRH, nearly 62 % RH at atmospheric standard condition). To
72 produce NH₄NO₃, nitrogen oxides (NO_x) and ammonia (NH₃) undergo a series of
73 chemical reactions. NO_x mostly emits as fresh NO, which is subsequently oxidized to
74 NO₂ and reacts with hydroxyl (OH) radicals to generate nitric acid (HNO₃), and then
75 HNO₃ reacts with NH₃ to yield NH₄NO₃ particles as listed in R1 and R2 (Calvert and
76 Stockwell, 1983). Particulate NH₄NO₃ formation rate is profoundly dependent on the
77 ambient T and RH since both parameters influence the equilibrium constant of NO₃⁻

78 and NH_4^+ between the particle and gas phases, as listed in R2 (Lin and Cheng, 2007).

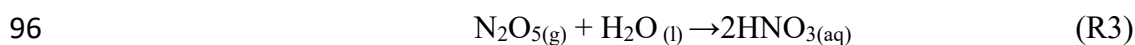


81 Here, k_1 and k_2 are the reaction rate and equilibrium constant of R1 and R2,

82 respectively. The equilibrium constant k_2 can be expressed as the product of HNO_3

83 and NH_3 .

84 Heterogeneous reactions have been considered an important mechanism of nitrate
85 formation during the nighttime. As listed in R3, liquid HNO_3 is produced by the
86 hydrolysis of dinitrogen pentoxide (N_2O_5) on aerosol surfaces (Brown & Stutz, 2012;
87 Chang et al., 2011; Mental et al., 1999; Wahner et al., 1998). Liquid HNO_3 can be
88 neutralized by NH_4^+ , which is produced from the conversion of gaseous NH_3 . Nitrate
89 aerosols yielded from both R2 and R3 require NH_3 , and we can therefore consider
90 these processes of NO_3^- formation to occur under NH_4^+ -rich conditions. Sometimes,
91 there is not enough NH_3 (NH_4^+) to react (to be neutralized) with HNO_3 (NO_3^-) after
92 complete neutralization by H_2SO_4 . Under this condition, HNO_3 tends to react (or to be
93 neutralized) with other alkaline species such as Ca-rich dust (CaCO_3), and
94 subsequently, nitrate aerosol is produced under a NH_4^+ -poor regime (Goodman et al.,
95 2000).



97

98 The Yangtze River Delta (YRD) region is one of the well-known polluted areas
99 in China (Zhang and Cao, 2015). Different from the case of dramatic elevated sulfate
100 aerosol levels in Beijing (Wang et al., 2016), nitrate aerosols seemed to be a major
101 contributing species during haze days in the YRD region (Wang et al., 2015; Wang et
102 al., 2018). The formation mechanisms of nitrate in Nanjing have not yet been well
103 understood, especially during high PM events. In this study, four intensive online

104 measurements of water-soluble ions in PM_{2.5} were conducted in Nanjing City in 2016
105 and 2017. The data provided information on the hourly evolution of water-soluble
106 inorganic ions (WSIIs) in the industrial city. The NO₃⁻ distributions under different
107 NH₄⁺ regimes (NH₄⁺-poor and NH₄⁺-rich conditions) were also discussed. Finally, we
108 investigated the potential formation mechanisms of nitrate aerosols and their
109 production rates during high-PM_{2.5} events based on the online measurements.

110

111 **2. Methodology**

112 **2.1 Sampling site**

113 Particulate WSIIIs and inorganic gases were continuously monitored at Nanjing
114 University of Information Science and Technology (NUIST) located in the northern
115 part of Nanjing City (see Figure S1). In addition to the contributions from vehicle
116 emissions, petroleum chemical refineries and steel manufacturing plants are situated
117 in the northeast and east direction at a distance of approximately 5 km. Four intensive
118 campaigns were conducted from March 2016 to August 2017. During each
119 experiment, the hourly concentrations of WSIIIs in PM_{2.5} and inorganic gases were
120 continuously observed. Meanwhile, the hourly PM_{2.5}, NO₂ and O₃ concentrations
121 along with the ambient T and RH were acquired from the Pukou air quality
122 monitoring station which is located to the southwest of the receptor site.

123

124 **2.2 Instruments**

125 To monitor the hourly concentrations of WSIIIs (Cl⁻, NO₃⁻, SO₄²⁻, Na⁺, NH₄⁺, K⁺,
126 Mg²⁺ and Ca²⁺), an online Monitor for Aerosols and Gases (MAGAR, Applikon-ENC,
127 The Netherlands) instrument with a PM_{2.5} inlet was employed. Using this instrument,
128 the WSIIIs in PM_{2.5} were collected by a stream jet aerosol collector, while acidic (HCl,
129 HONO, HNO₃ and SO₂) and basic gases (NH₃) were dissolved in a hydrogen peroxide

130 solution on a wet rotation denuder (ten Brink et al., 2007; Griffith, et al., 2015). The
131 liquid samples were then collected with syringe pumps and analyzed by ion
132 chromatography (IC). Before each campaign, a seven-point calibration curve of each
133 species was made, and an internal standard solution (LiBr) was used to check
134 instrumental drifts. The method detection limits (MDLs) of Cl^- , NO_3^- , SO_4^{2-} , Na^+ ,
135 NH_4^+ , K^+ , Mg^{2+} and Ca^{2+} were, 0.01, 0.04, 0.06, 0.05, 0.05, 0.07, 0.05 and $0.11 \mu\text{g m}^{-3}$
136 3 , respectively. For gases, the MDLs were 0.07, 0.09, 0.06, 0.02 and $0.08 \mu\text{g m}^{-3}$ for
137 HCl, HONO, HNO_3 , SO_2 and NH_3 , respectively.

138

139 **2.3 ISORROPIA-II model**

140 In this work, we used the ISORROPIA-II model to calculate the aerosol liquid
141 water content (ALWC). ISORROPIA II is a thermodynamic equilibrium model which
142 is built based on the $\text{Na}^+ - \text{Cl}^- - \text{Ca}^{2+} - \text{K}^+ - \text{Mg}^{2+} - \text{SO}_4^{2-} - \text{NH}_4^+ - \text{NO}_3^- - \text{H}_2\text{O}$ aerosol
143 system (Fountoukis & Nenes, 2007). This model has been successfully used to
144 estimate the liquid water content in aerosols with uncertainty of $\sim 20\%$ compared to
145 the observed ALWC (Bian et al., 2014; Guo et al., 2015; Liu et al., 2017). This
146 underestimation might be due to the missed species in ISORROPIA II, organic
147 aerosols, , which contributed approximately 27% of the total ALWC (Bougiatioti et
148 al., 2007). Here, the model was computed as a “forward problem”, in which the
149 quantities of aerosol- and gas-phase compositions along with the T and RH were well
150 known. Additionally, the modeled values were determined using the “metastable”
151 mode, which indicated that the aerosol compositions were assumed to be composed of
152 an aqueous solution (Liu et al., 2017). The details of this model can be found
153 elsewhere (Fountoukis and Nenes, 2007). In this work, the observed concentrations of
154 total nitrate ($\text{HNO}_3 + \text{NO}_3^-$), total ammonium ($\text{NH}_3 + \text{NH}_4^+$), total chloride ($\text{HCl} + \text{Cl}^-$),
155 SO_4^{2-} , Na^+ , K^+ , Mg^{2+} and Ca^{2+} along with measured ambient T and RH served as

156 input of ISORROPIA II model.

157

158 **2.4 Potential source contribution function**

159 Potential source contribution function (PSCF) is a method to identify the
160 potential source regions of air pollutants. It has also been widely used to differentiate
161 local emission from long-range transported pollution (Zhang et al., 2013; Hui et al.,
162 2018) based on the trajectory analysis calculated from GDAS (Global Data
163 Assimilation System), which processed by the National Centers for Environmental
164 Prediction (NCEP). The zone of concern is divided into $i \times j$ small equal grid cells and
165 then PSCF in the i - j th cell ($PSCF_{ij}$) can be defined as (Polissar et al., 1999):

$$166 \quad PSCF_{ij} = \frac{m_{ij}}{n_{ij}} \quad (1)$$

167 where m_{ij} is the number of “high nitrate pollution” trajectory endpoints in the i - j th cell
168 and n_{ij} is the total number of trajectory endpoints fallen into the i - j th cell. In this
169 study, the 80th percentile value of nitrate concentration was treated as “high nitrate
170 pollution” threshold. To reduce the uncertainty caused by the small values of n_{ij} , the
171 weighting function of W_{ij} has to be considered (Polissar et al., 1999):

172

$$173 \quad W_{ij} = \left\{ \begin{array}{l} 1.00; 80 < n_{ij} \\ 0.70; 20 < n_{ij} \leq 80 \\ 0.42; 10 < n_{ij} \leq 20 \\ 0.05; n_{ij} \leq 10 \end{array} \right.$$

177

178 In this study, the domain of the study area was in a range of 20-55 °N and 105-135 °E;
179 the resolution of grid cell was 0.5°×0.5°.

180

181 3. Results and discussion

182 3.1 Overview of water-soluble inorganic ions

183 Four intensive online measurements of WSIs in PM_{2.5} were carried out in
184 Nanjing City from March 2016 to August 2017. Figure 1a plots the time series of the
185 hourly PM_{2.5} mass concentrations during the sampling periods. As seen, the hourly
186 PM_{2.5} mass concentrations varied from 5 to 252 $\mu\text{g m}^{-3}$ with a mean value of 58 ± 35
187 $\mu\text{g m}^{-3}$. Compared with the 24-hour guideline (25 $\mu\text{g m}^{-3}$) suggested by the World
188 Health Organization (WHO), our average PM_{2.5} concentration ($58 \mu\text{g m}^{-3}$) was 2.3
189 times higher. This indicated that PM pollution in Nanjing City was a serious problem.
190 During the campaigns, several high-PM_{2.5} events with hourly PM_{2.5} concentrations of
191 higher than 150 $\mu\text{g m}^{-3}$ were observed in the springtime and wintertime. These high
192 PM_{2.5} levels lasted for more than 3 hours, with obviously elevated NO₃⁻. The details
193 of nitrate formation during the high-PM_{2.5} hours will be discussed in the following
194 sections.

195 Figure 1b shows the time series of the hourly concentrations of SIA species,
196 including SO₄²⁻, NO₃⁻ and NH₄⁺. The lack of data from March 7 to 14, 2016 was due
197 to a malfunction of the MARGA instrument. During the sampling periods, the NO₃⁻
198 concentrations varied from 0.1 to 85.1 $\mu\text{g m}^{-3}$ with a mean value of $16.7 \pm 12.8 \mu\text{g m}^{-3}$.
199 The SO₄²⁻ concentrations ranged from 1.7 to 96.2 $\mu\text{g m}^{-3}$ and averaged 14.9 ± 9.1
200 $\mu\text{g m}^{-3}$. The NH₄⁺ concentrations fluctuated between 0.8 and 44.9 $\mu\text{g m}^{-3}$ with a mean
201 value of $10.7 \pm 6.7 \mu\text{g m}^{-3}$. On average, SIA accounted for 91 % of the total water-
202 soluble inorganic ions (TWSIs) during the entirety of the sampling periods (see
203 Figure 2a). Among these species, NO₃⁻ accounted for 35 % of the TWSIs, followed
204 by SO₄²⁻ (33 %) and NH₄⁺ (24 %). The abundances of other ions, including Cl⁻, K⁺,
205 Ca²⁺, Na⁺ and Mg²⁺, were 5, 2, 1, 0.7 and 0.3 %, respectively. Figure S2 shows the
206 scatter plot of the equivalent concentrations of the cations (Na⁺, NH₄⁺, K⁺, Mg²⁺ and

207 Ca^{2+}) and anions (Cl^- , SO_4^{2-} and NH_4^+). As seen, good correlations ($R = 0.98 - 0.99$,
208 with a significance level $p < 0.05$) between cations and anions were found during the
209 various sampling periods. The ratio of cation-to-anion was very close to 1.0 during
210 each season, reflecting an ionic balance. This also indicated that our data exhibited
211 good quality and was able to be used for the further analysis of scientific issues.

212 All SIA species exhibited similar seasonal patterns, with lower concentrations in
213 the summer, especially for NO_3^- . The average concentrations of nitrate were 6.7 and
214 $5.7 \mu\text{g m}^{-3}$ in the summertime of 2016 and 2017, respectively (see Figure S3). These
215 values were much lower than those observed during other seasons. The local
216 meteorological conditions, which were favorable for the dilution of air pollution, were
217 one of the reasons for the declined NO_3^- concentrations during the hot seasons (Zhang
218 and Cao, 2015). Another important reason for this effect was attributed to the
219 formation process of $\text{PM}_{2.5}$ nitrate, which is very sensitive to the ambient T and RH
220 (Lin and Cheng, 2007). Figure S4a depicts the theoretical equilibrium constants of
221 partitioned NO_3^- and NH_4^+ between the particle and gas phase ($\text{HNO}_{3(\text{g})} + \text{NH}_{3(\text{g})} \rightarrow$
222 $\text{NH}_4\text{NO}_{3(\text{s, aq})}$ as seen in R2) under different T and RH conditions. The details of
223 calculation approach of the theoretical equilibrium constants are described in
224 *Supplementary S1*. Note that the *Y-axis* is presented on a log scale. The theoretical
225 equilibrium constants increased exponentially with increasing ambient temperature
226 but decreased with increasing RH. This indicated that NH_4NO_3 would be partitioned
227 into the gas phase due to high equilibrium constants under high-temperature and low-
228 RH conditions. Figure S4b illustrates the time series of the theoretical and observed
229 equilibrium constants during the sampling periods. As can be seen, most of the
230 observed equilibrium constants were higher than the theoretical ones, suggesting that
231 NH_4NO_3 aerosols were produced in Nanjing during the sampling periods. Obviously,
232 higher theoretical and lower observed equilibrium constants were found during the

233 summer. This suggested that more NO_3^- and NH_4^+ would tend to be partitioned into
234 the gas phase, resulting in lower particulate nitrate concentrations during hot seasons
235 (Lin and Cheng, 2007).

236 Apart from seasonal variations, pronounced diurnal patterns were also found for
237 SIA species (see Figure 3). NO_3^- exhibited similar diel cycles during different seasons,
238 with higher concentrations in the early morning (3 a.m. - 7 a.m.) and lower levels
239 between 2 p.m. and 5 p.m. The high nitrate concentrations in the early morning might
240 be caused by the nitrate formation via heterogeneous reaction in the dark, and gas-
241 phase oxidation after sunrise and the subsequent condensation on pre-existing
242 particles before the temperature increased and RH decreased afterwards. Moreover,
243 the lower planet boundary layer (PBL) might be another reason for enhanced nitrate
244 in the early morning. However, the lower concentrations of nitrate during the daytime
245 might be attributed to the higher PBL, and high temperatures, which inhibited the
246 build-up of nitrate, especially during the summertime. In terms of sulfate, higher
247 concentrations were observed between 6 am. and 1 p.m., indicating that the formation
248 rate of sulfate was higher than the removal/dilution rate, leading to an increase of the
249 sulfate concentration during the daytime. The diurnal patterns of NH_4^+ mimicked
250 those of NO_3^- , showing lower concentrations during the daytime. This was explained
251 by the drastic decrease of particulate NH_4NO_3 concentrations under high temperatures
252 and low relative humidity, resulting in lower NH_4^+ levels during the daytime.

253

254 **3.2 Enhancements of nitrate at high $\text{PM}_{2.5}$ levels**

255 Figure S5 shows the scatter plots of NO_3^- , SO_4^{2-} and NH_4^+ against $\text{PM}_{2.5}$. As
256 seen, the slopes of NO_3^- (NO_3^- vs. $\text{PM}_{2.5}$ mass), SO_4^{2-} and NH_4^+ were 0.30, 0.24 and
257 0.19, respectively. This suggested that the increasing rate of NO_3^- during the high-
258 $\text{PM}_{2.5}$ events was higher than those of other SIA species. At high $\text{PM}_{2.5}$ levels ($\text{PM}_{2.5} \geq$

259 150 $\mu\text{g}/\text{m}^3$), NO_3^- , SO_4^{2-} and NH_4^+ contributed 39, 28 and 24 % of the TWSIIs,
260 respectively (Figure 2b). However, the relative abundances of NO_3^- , SO_4^{2-} and NH_4^+
261 during low $\text{PM}_{2.5}$ concentrations (hourly $\text{PM}_{2.5} < 35 \mu\text{g}/\text{m}^3$, see Figure 2c) were 29, 37
262 and 23 %, respectively. In recent years, dramatically enhanced amounts of nitrate
263 aerosols during high-PM events have been observed at many urban sites in China
264 (Wen et al., 2015; Wang et al., 2017; 2018; Zou et al., 2018). For instance, Zou et al.
265 (2018) found that the nitrate concentrations during the occurrence of polluted air in
266 Beijing and Tianjin were almost 14 times higher than those on relatively clean days
267 ($\text{PM}_{2.5} < 75 \mu\text{g}/\text{m}^3$), and the enhancement ratio of nitrate was much higher than that
268 (5.3) of sulfate. Wang et al. (2018) noted that the enhancement ratio of NO_3^- (~6)
269 between haze and clear days in Ningbo of the YRD region was much higher than that
270 of SO_4^{2-} (~3). These findings suggested that NO_3^- was a major contributing species to
271 fine particles during haze days since its increasing ratio between haze and non-haze
272 days was much higher than those of other SIA species, such as sulfate and
273 ammonium.

274

275 **3.3 PSCF result of high nitrate pollution**

276 During the high $\text{PM}_{2.5}$ pollution, significant enhanced nitrate aerosols in terms
277 of both absolute concentration and relative abundance to TWSIIs were found. Next,
278 we tried to use PSCF analysis to identify whether local emission or long-range
279 transported pollution was the major source of high nitrate concentrations at the
280 receptor site. In this work, the 80th percentile values of nitrate concentration was
281 selected as “high nitrate pollution” threshold for PSCF analysis. Figure 4 plots the
282 PSCF result of high nitrate pollution in Nanjing during the sampling periods. The
283 region corresponding to high PSCF value grid is a potential source region of nitrate
284 aerosols. As can be seen, the areas with high PSCF value (>0.8) were regularly local

285 areas surrounding by Nanjing while PSCF values from other long-distance areas were
286 lower than 0.2. This suggested that NO_3^- aerosols in Nanjing during the high nitrate
287 pollution were likely from local emissions rather than long-range transported sources.

288

289 **3.4 Nitrate formation under different ammonium regimes**

290 Ammonium is a major species that neutralizes particulate SO_4^{2-} and NO_3^- . In the
291 atmosphere, SO_4^{2-} competes with NO_3^- for NH_4^+ during their formation processes, and
292 therefore, the relationship between the molar ratios of $\text{NO}_3^-/\text{SO}_4^{2-}$ and $\text{NH}_4^+/\text{SO}_4^{2-}$ can
293 give us a hint for understanding the formation of NO_3^- under different ammonium
294 regimes (Pathak et al., 2009; He et al., 2012; Tao et al., 2016). In an ammonium-rich
295 regime, the HNO_3 produced by both gas oxidation and heterogeneous process reacts
296 (or neutralizes) with “excess-ammonium” (excess- NH_4^+) at a $\text{NH}_4^+/\text{SO}_4^{2-}$ molar ratio
297 > 2 (theoretical value in an NH_4^+ -rich regime) when sulfate is completely neutralized
298 by NH_4^+ to form $(\text{NH}_4)_2\text{SO}_4$ (Squizzato et al., 2013; Ye et al., 2011). In contrast,
299 nitrate can be found under ammonium-poor conditions with a theoretical $\text{NH}_4^+/\text{SO}_4^{2-}$
300 value that should be less than 2 (Pathak et al., 2009). Under NH_4^+ -poor conditions,
301 HNO_3 reacts with other cations, such as the calcium carbonate frequently found in
302 natural dust.

303 Figure 5 shows the scatter plot of the molar ratios of $\text{NO}_3^-/\text{SO}_4^{2-}$ against
304 $\text{NH}_4^+/\text{SO}_4^{2-}$. It is found that good correlations exist between $\text{NO}_3^-/\text{SO}_4^{2-}$ and
305 $\text{NH}_4^+/\text{SO}_4^{2-}$ under NH_4^+ -rich regimes, with a coefficient of determination (R^2) of 0.84
306 - 0.94 in the different seasons (see in Table 1). Utilizing the linear regression model,
307 we suggested that nitrate aerosols (in NH_4^+ -rich regimes) began to form when the
308 $\text{NH}_4^+/\text{SO}_4^{2-}$ molar ratios exceeded the criterion values of 1.7-2.0 during the different
309 seasons (Table 1). The criterion value can be calculated as absolute value of
310 “intercept” dividing by slope in each linear regression model (He et al., 2012). The

311 criterion values below 2 suggested that part of the sulfate might have existed in other
312 forms, such as ammonium bisulfate. On the other hand, under ammonium-rich
313 conditions, nitrate concentrations should be positively proportional to “excess-NH₄⁺”
314 concentrations, a relationship which was defined as [excess-NH₄⁺] = (NH₄⁺/SO₄²⁻ -
315 criterion value) × [SO₄²⁻] (Pathak et al., 2009) (sulfate is in the units of nmol m⁻³
316 here). The criterion values were acquired from the regression models, as listed in
317 Table 1. The results revealed that the excess-NH₄⁺ concentrations varied from -283 to
318 1422 nmol m⁻³ (see Figure 6), and only 1 % of data showed deficit-NH₄⁺ conditions,
319 reflecting that NO₃⁻ formation in Nanjing occurred primarily under the NH₄⁺-rich
320 conditions. Moreover, the excess-NH₄⁺ had apparent diurnal cycles, with higher
321 concentrations in the early morning and lower concentrations at midday and in the
322 early afternoon (see Figure 3, where we converted the units from nmol m⁻³ to μg m⁻³).
323 The diurnal patterns of NO₃⁻ mimicked those of the excess-NH₄⁺. This also suggested
324 that particulate NO₃⁻ formation occurred mainly under NH₄⁺-rich conditions. Figure 6
325 illustrates the relationship between the nitrate and excess-NH₄⁺ molar concentrations
326 during the sampling periods. The nitrate molar concentrations correlated linearly with
327 the excess-NH₄⁺ molar concentrations with a slope of approximately 1.0, which was
328 consistent with the molar ratio of reaction between HNO₃ and NH₃. Interestingly,
329 some scattered points were found in high ammonium concentrations (excess-NH₄⁺ ≥
330 900 nmol m⁻³ ~ 16.2 μg m⁻³), implying that residual NH₄⁺ might be presented in
331 another form such as NH₄Cl under high-NH₄⁺ conditions. On the contrary, NO₃⁻
332 aerosols can be produced without involving NH₃; therefore, NO₃⁻ did not correlate
333 well with the excess NH₄⁺ under a NH₄⁺-poor regime.

334 In this study, high nitrate concentrations were always found under NH₄⁺-rich
335 regimes, elucidating that nitrate production during high PM levels in Nanjing had to
336 be involved with NH₃ or NH₄⁺. Figure 6 also shows the nitrate concentrations against

337 the excess-NH₄⁺ observed in various cities of China during the summertime (Pathak et
338 al., 2009; Griffith et al., 2015). In Beijing and Shanghai, high nitrate concentrations
339 during the summertime were found under NH₄⁺-deficient conditions, which was very
340 different from the findings of this work. In these studies (Pathak et al., 2009; Griffith
341 et al., 2015), the high nitrate concentrations associated with NH₄⁺-poor conditions
342 might be due to the lower excess-NH₄⁺ concentrations under high-SO₄²⁻ conditions at
343 that time since the strict control of SO₂ emissions by the Chinese government started
344 in 2010 (Zheng et al., 2018). In recent years, the reduction of anthropogenic SO₂
345 emissions decreased the airborne SO₄²⁻ concentrations, resulting in more excess-NH₄⁺
346 and leading to nitrate aerosol formation under NH₄⁺-rich regimes. This argument can
347 be supported by the recent results shown in Figure S6, in which high nitrate
348 concentrations in Beijing were always found under NH₄⁺-rich regimes.

349

350 **3.5 Nitrate formation mechanism during high-PM_{2.5} episodes**

351 In this section, we attempted to explore the formation mechanisms of nitrate
352 aerosols during high PM_{2.5} levels. Here, nitrogen conversion ratio (Fn) was used to
353 evaluate the conversion capability of NO₂ to total nitrate (TN, TN=HNO₃ + NO₃⁻),
354 and it can be defined as (Khoder, 2002; Lin et al., 2006):

$$355 \quad F_n = \frac{GNO_3^- + PNO_3^-}{GNO_3^- + PNO_3^- + NO_2} \quad (1)$$

356

357 where GNO₃⁻ and PNO₃⁻ represent the NO₂ concentrations in nitric acid and
358 particulate nitrate, respectively, with the units of μg m⁻³. The results showed that the
359 Fn values during the sampling periods varied from 0.01 to 0.57 with a mean value of
360 0.14 ± 0.09 (see Figure 1e). This value was comparable to that (0.17) in Taichung,
361 Taiwan, where both gas-oxidation and heterogeneous reaction were the dominant
362 formation mechanisms of atmospheric HNO₃ (or NO₃⁻) (Lin et al., 2006). However,

363 our F_n value was 2.3 time higher than that (0.06) in Dokki, Egypt (Khoder, 2002).
364 The reason of significant discrepancy of F_n between this work and that in Dokki was
365 not clearly understood, but it might be attributed to different formation processes of
366 HNO_3 . In Dokki, gas-phase oxidation was the dominant pathway of HNO_3 production
367 while heterogeneous process (R3) played an important role in HNO_3 formation in
368 addition to gas-phase oxidation in Nanjing, especially during the high- $\text{PM}_{2.5}$ events
369 (discussed later). The reaction rate of HNO_3 by heterogeneous process was much
370 higher than that by gas-phase oxidation (Calvert and Stockwell, 1983) and therefore,
371 the F_n value was much higher in this study. On the other hand, F_n displayed
372 significant diurnal cycles, with the highest value in the early morning (see in Figure
373 3). This elevated F_n coincided with increasing ALWC, suggesting heterogeneous
374 reaction since ALWC is one of the key parameters which favors the transformation of
375 N_2O_5 to liquid HNO_3 in this process (also indicated that nitrate formation was
376 associated with heterogeneous process). On the contrary, a second peak of F_n was
377 found in the early afternoon when O_x ($\text{O}_x = \text{NO}_2 + \text{O}_3$, an index of the oxidation
378 capacity) concentrations increased, but ALWC decreased. This suggested that the
379 HNO_3 formation might be mainly associated with the gas-phase reaction of $\text{NO}_2 +$
380 OH during the daytime; also reflected that nitrate formation was via gas-phase
381 oxidation.

382 Assuming that long-range transported nitrate can be neglected in this study (in
383 section 3.3), we attempted to analyze the correlations of F_n vs. OH and F_n vs. ALWC
384 to investigate whether gas-phase oxidation or heterogeneous reactions might be the
385 dominant mechanism of nitrate production. In this work, the OH radical
386 concentrations were not measured; hence, we used O_x as a proxy of OH . The ALWC
387 was acquired by computing the ISOPROPIA II model as described in section 2.3.
388 Figure 7 illustrates the scatter plots of F_n against O_x and ALWC in both daytime and

389 nighttime aerosol samples during the high-PM_{2.5} events. Fn correlated well with the
390 ALWC, with a correlation coefficient (R) of 0.72 and 0.76 ($p < 0.05$) at daytime and
391 nighttime samples, respectively. However, a poor correlation was found between Fn
392 and Ox (R was 0.17 and 0.52 for the daytime and nighttime samples, $p > 0.05$). This
393 implied that nitrate formation during the high-PM_{2.5} events in Nanjing was likely
394 attributed to heterogeneous reactions. This result was consistent with recent
395 conclusions reached by oxygen isotope techniques, in which the hydrolysis of N₂O₅ in
396 preexisting aerosols was found to be a major mechanism of NO₃⁻ formation (Chang et
397 al., 2018).

398

399 **3.6 Case study and production rate of NO₃⁻ during PM_{2.5} episodes**

400 Figure 8 shows several high-PM_{2.5} events observed from March 3 to 6 in 2016.
401 In case I, the high PM_{2.5} concentrations started at 6 p.m. on March 3 and ended at 3
402 a.m. on March 4. During this event, the SO₄²⁻ and NH₄⁺ concentrations remained at
403 almost constant levels, but the NO₃⁻ concentrations revealed a slight enhancement. In
404 the early morning of March 4, the NO₃⁻ concentrations increased from 39.4 to 47.8 μg
405 m⁻³ within 4 hours, resulting in a nitrate production rate of 2.3 μg m⁻³ h⁻¹ (~5.5 % h⁻¹,
406 the calculation of NO₃⁻ production rate can be seen in the *Supplementary S2*). In case
407 II, high PM_{2.5} concentrations were observed from 8. a.m. to 2. p.m. on March 4. The
408 NO₃⁻ concentrations were much higher than those of SO₄²⁻, indicating nitrate-
409 dominated aerosols. In this case, the NO₃⁻ concentrations increased from 38.1 to 51.2
410 μg m⁻³ within 6 hours, suggesting that the increasing rate of NO₃⁻ was 1.0 μg m⁻³ h⁻¹
411 (2.4 % h⁻¹). Since the high NO₃⁻ concentrations occurred under high-Ox and low-
412 ALWC conditions, this suggested that the gas-phase reaction of NO₂ + OH might be
413 the dominant source of NO₃⁻ production in this event. In case III, a rapid growth of the
414 PM_{2.5} mass was found around midnight, along with a dramatic increase of NO₃⁻

415 concentrations from 11 p.m. on March 4 ($31.0 \mu\text{g m}^{-3}$) and maximizing at 1 a.m. the
416 next day ($64.5 \mu\text{g m}^{-3}$). The increasing rate of NO_3^- was estimated to be $11.4 \mu\text{g m}^{-3} \text{ h}^{-1}$
417 ($\sim 19.5 \% \text{ h}^{-1}$), which was much higher than those in case I and II. The high-nitrate
418 event was found under increasing ALWC and decreasing Ox concentration conditions,
419 suggesting that nitrate production occurred through heterogeneous processes. In case
420 IV, the enhancements of all SIA species coincided with increasing ALWC and
421 declining Ox concentrations. Again, the enhancement of nitrate was likely attributed
422 to heterogeneous reactions rather than to gas-phase processes. In these events, the
423 NO_3^- production rate was estimated to be $5.0 \mu\text{g m}^{-3} \text{ h}^{-1}$ ($\sim 15.4 \% \text{ h}^{-1}$).
424 Through the sampling periods, a total of twelve high $\text{PM}_{2.5}$ events were found, and the
425 NO_3^- concentrations increased significantly during all the episodes (see in Table S1).
426 Seven episodes suggested that heterogeneous processes ($\text{N}_2\text{O}_5 + \text{H}_2\text{O}$) might be a
427 major pathway for nitrate formation since elevated NO_3^- levels coincided with
428 increasing ALWC and decreasing Ox (or Ox remaining at a constant level). Among
429 these heterogeneous process events, five cases (Case III, Case IX, Case X, Case XI
430 and Case XII in Table S1) were observed during the nighttime (5 p.m. – 6 a.m. on the
431 next day). This suggested that approximately 70 % heterogeneous reaction of nitrate
432 production was observed in the dark. In these events, the average NO_3^- growth rate
433 was $12.6 \pm 7.3 \% \text{ h}^{-1}$ ($4.1 \pm 3.6 \mu\text{g m}^{-3} \text{ h}^{-1}$). This value was in agreement with those in
434 the literatures which the production rate of nitrate via heterogeneous reaction were
435 $14.3 \% \text{ h}^{-1}$ by both field measurements and laboratory works (Calvert and Stockwell,
436 1983; Pathak et al., 2011). On the contrary, NO_3^- concentrations rose with increasing
437 Ox and decreasing ALWC in two $\text{PM}_{2.5}$ episodes, indicating gas-phase processes
438 ($\text{NO}_2 + \text{OH}$). As listed in Table S1, these gas-phase reaction cases occurred mainly
439 during the daytime. The average production rate of NO_3^- in the gas-oxidation reaction
440 cases averaged $2.5 \pm 0.1 \% \text{ h}^{-1}$ ($0.8 \pm 0.3 \mu\text{g m}^{-3} \text{ h}^{-1}$), which was in line with that (2.4

441 % h⁻¹) in the subtropical polluted urban site that nitrate aerosols were mainly from
442 gas-oxidation process (Lin et al., 2007). Moreover, we also found some cases in
443 which the elevated NO₃⁻ might have been from both gas-phase and heterogeneous
444 reactions, and the corresponding NO₃⁻ growth rate was approximately 7.5 ± 3.0 % h⁻¹
445 (2.5 ± 0.2 μg m⁻³ h⁻¹). In conclusion, enhancements of NO₃⁻ in Nanjing usually
446 occurred under increased ALWC and decreased Ox conditions, indicating that
447 heterogeneous reactions provided the dominant pathway of nitrate formation during
448 the PM_{2.5} episodes. Moreover, the average growth rate of NO₃⁻ (12.6 % h⁻¹) by
449 heterogeneous processes was 5 times higher than that (2.5 % h⁻¹) of gas-phase
450 reactions. This might explain the abrupt increase of nitrate concentrations during the
451 high PM_{2.5} events.

452

453 **3.7 HNO₃/NH₃ limitation of nitrate aerosol formation**

454 In Nanjing, high nitrate concentrations occurred mainly under NH₄⁺-rich
455 regimes, indicating the involvement of atmospheric NH₃. This also demonstrated that
456 both HNO₃ and NH₃ were crucial precursors for particulate nitrate formation. In this
457 section, we attempted to discuss whether HNO₃ or NH₃ was the limited factor for
458 nitrate formation in Nanjing during the high-PM_{2.5} events. ISORROPIA II model is
459 capable of predicting concentrations of particulate ions in addition to ALWC under
460 thermodynamic equilibrium between gas- and aerosol-phase of these ions (Tang et al.,
461 2016). In section 3.5, we used this model to estimate ALWC. Indeed, the output data
462 also included concentrations of ionic species. Figure S7 illustrates the scatter plots of
463 modeled results against observations of NO₃⁻, SO₄²⁻ and NH₄⁺ in Nanjing during the
464 sampling periods. Good correlations were found between modeled results and
465 observations (R²=0.97-0.99 with all slopes of approximately 1.0), suggesting that
466 ISORROPIA II had a good performance in prediction of SIA species. As a result, we

467 can use ISORROPIA II model to test sensitivity of HNO₃ and NH₃ to particulate
468 nitrate concentrations (Guo et al., 2018).

469 Figure 7 shows the contour plot of the simulated nitrate concentrations depending
470 on the various total nitrate (TN) and total ammonium (TA, TA=NH₃ + NH₄⁺) levels
471 under thermodynamic equilibrium conditions computed by ISORROPIA II model.
472 The details of considered chemical reactions in ISORROPIA II model can be seen
473 elsewhere (Fountoukis & Nenes, 2007). Here, sulfate concentrations were assumed to
474 be 10 and 60 μg m⁻³ for the tests of different sulfate conditions. The average
475 concentrations of total chloride (HCl + Cl⁻, 1.3 μg m⁻³), Na⁺ (0.2 μg m⁻³), K⁺ (0.8 μg
476 m⁻³), Mg²⁺ (0.1 μg m⁻³) and Ca²⁺ (0.5 μg m⁻³) along with ambient T (20 °C) and RH
477 (62 %) at the receptor site during the sampling period served as input data in this
478 model. The results showed that the lower simulated NO₃⁻ concentrations was found in
479 the higher SO₄²⁻ case. This was attributed to less NH₄NO₃ formation under higher
480 SO₄²⁻ conditions since SO₄²⁻ would compete with NO₃⁻ for NH₄⁺.

481 According to the simulated results, we can roughly split the plots into two parts:
482 one is HNO₃-limited area (right), and another is NH₃-limited region (left). The
483 observed TN and TA concentrations (pink circles) in Nanjing are also plotted in this
484 figure. Most of the observed data sets were mainly affected by TN under a low-SO₄²⁻
485 case. Under a high-SO₄²⁻ condition, the observed data fell into TA-limited under a
486 low-TN and -TA regime, but fell into TN-limited in high-TA and-TN regimes. During
487 the sampling period, high nitrate concentrations always accompanied with high TN
488 and TA levels, highlighting that nitrate aerosol production in Nanjing during the high
489 PM_{2.5} levels was mainly control by HNO₃. Therefore, control of NO_x emissions,
490 which reduced HNO₃ concentrations, might be an important way to decrease airborne
491 nitrate concentrations and ameliorate the air quality in Nanjing.

492

493 4. Conclusion and remarks

494 Four intensive online measurements of water-soluble ions in PM_{2.5} were carried
495 out in Nanjing City in 2016 and 2017 to realize the evolutions of SIA and the potential
496 formation mechanisms of particulate nitrate. During the sampling periods, the average
497 concentrations of NO₃⁻, SO₄²⁻ and NH₄⁺ were 16.7, 14.9 and 10.7 μg m⁻³, respectively.
498 This indicated that NO₃⁻ dominated the SIA. Significant seasonal variations and
499 diurnal cycles were found for all SIA species. The low NO₃⁻ concentrations observed
500 during the summer daytime could be attributed to the enhanced theoretical and
501 declined observed equilibrium constants of NO₃⁻ and NH₄⁺ between gas- and particle-
502 phase. Obvious enhancements of NO₃⁻ were found in terms of both absolute
503 concentrations and relative abundances during the PM_{2.5} episodes, indicating that
504 NO₃⁻ was a major contributing species to PM_{2.5}. Different from the results obtained in
505 Beijing and Shanghai, high nitrate concentrations always occurred under NH₄⁺-rich
506 regimes. The nitrogen conversion ratio, Fn, correlated well with the ALWC but not
507 with Ox during high-PM_{2.5} episodes. These findings indicated that NO₃⁻ aerosols at
508 the receptor site were mainly produced by heterogeneous reactions (N₂O₅ + H₂O) with
509 the involvement of NH₃. The average production rate of NO₃⁻ from heterogeneous
510 reactions was estimated to be 12.6 % h⁻¹, which was 5 time higher than that of gas-
511 phase reactions. According to the observations and ISORROPIA II simulated results,
512 particulate nitrate formation in Nanjing was HNO₃-limited, suggesting that the control
513 of NO_x emissions will be able to decrease the nitrate concentration and improve the
514 air quality in this industrial city.

515 During the last decade, the mass ratios of nitrate-to-sulfate in PM_{2.5} in the YRD
516 region have been found to range from 0.3 to 0.7 (Lai et al., 2007; Wang et al., 2003;
517 2006; Yang et al., 2005; Yao et al., 2002), reflecting that the SO₄²⁻ concentration was
518 much higher than the NO₃⁻ concentration. In the current study, the average mass ratio

519 of nitrate-to-sulfate was 1.1. Indeed, high nitrate-to-sulfate mass ratios of > 1 were
520 also observed in other mega-cities of China recently (Ge et al., 2017; Wei et al., 2018;
521 Ye et al., 2017; Zou et al., 2018). The elevated nitrate-to-sulfate ratio should be due to
522 the dramatic reduction of SO₂ emissions. The enhanced ratio also suggests that we
523 should pay more attention to and develop some strategies for the reduction of NO_x
524 emissions, leading to declined nitrate concentrations in the atmosphere and
525 improvement of the air quality in China.

526

527 **Data availability**

528 All the data used in this paper are available from the corresponding author upon
529 request (dryanlinzhang@outlook.com or zhangyanlin@nuist.edu.cn).

530

531 **Author contributions**

532 YLZ conceived and designed the study. YCL analyzed the data and wrote the
533 manuscript with YLZ. FM and MB performed aerosol sampling and data analyses
534 with YCL.

535

536 **Competing interests**

537 The authors declare that they have no conflict of interest.

538

539 **Acknowledgements**

540 This study was financially supported by the National Key R&D Program of China
541 (Grant No. 2017YFC0212700), the Natural Scientific Foundation of China (Nos.
542 41761144056, 91644103 and 41977185) and Jiangsu Innovation & Entrepreneurship
543 Team.

544

545 **References**

- 546 Baasandorj, M., Hoch, S. W., Bares, R., Lin, J. C., Brown, S. S., Millet, D. B., Martin,
547 R., Kelly, K., Zarzana, K. J., Whiteman, C. D., Bube, W. P., Tonnesen, G.,
548 Jaramillo, J. C. and Sohl, J.: Coupling between chemical and meteorological
549 processes under persistent coal-air poor conditions: evolution of wintertime PM_{2.5}
550 events and N₂O₅ observation in Utah's Salt Lake Valley. *Environ. Sci. Technol.*,
551 **51**, 5941-5950, <https://doi.org/10.1021/acs.est.6b06603>, 2017.
- 552 Bian, Y. X., Zhao, C. S., Ma, N., Chen, J., and Xu, W. Y.: A study of aerosol liquid
553 water content based on hygroscopicity measurements at high relative humidity in
554 the Northern China Plain. *Atmos. Chem. Phys.*, **14**, 6417-6426,
555 <https://doi.org/10.5194/acp-14-6417-2014>, 2014.
- 556 Bougiatioti, A., Nikolaou, P., Stavroulas, I., Kouvarakis, G., Weber, A., Nenes, R.,
557 Kanakidou, M., and Mihalopoulos, N.: Particle water and pH in the eastern
558 Mediterranean: source variability and implication of nutrient availability. *Atmos.*
559 *Chem. Phys.*, **16**, 4579-4591, <https://doi.org/10.5194/acp-7-4639-2007>, 2007.
- 560 Brauer, M., Hoek, G., Vliet, V. P., Meliefste, K., Fischer, P. H., Wijga, A., Koopman,
561 L. P., Neijens, H. J., Gerritsen, J., Kerkhof, M., Heinrich, J., Bellander, T., and
562 Brunekreef, B.: Air pollution from traffic and the development of respiratory
563 infections and asthmatic and allergic symptoms in children. *Am. J. Respir. Crit.*
564 *Care Med.*, **166**, 1092-1098, <https://doi.org/10.1146/rccm.200108-007OC>, 2002.
- 565 Brown, S. S., and Stutz, J.: Nighttime radical observation and chemistry. *Chem. Soc.*
566 *Rev.*, **41**, 6405-6447, <https://doi.org/10.1039/c2cs35181a>, 2012.
- 567 Calvert, J. G., and Stockwell, W. R.: Acid generation in the troposphere by gas-phase
568 chemistry. *Environ. Sci. Technol.*, **17**, 428-443,
569 <https://doi.org/10.1021/es00115a727>, 1983.
- 570 Chan, C. K., and Yao, X.: Air pollution in mega cities in China. *Atmos. Environ.*, **42**, 1-

571 42, <https://doi.org/10.1016/j.atmosenv.2007.09.003>, 2008.

572 Chang, W. L., Bhave, P. V., Brown, S. S., Riemer, N., Stutz, J., and Dabdub, D.:
573 Heterogeneous atmospheric chemistry, ambient measurements, and model
574 calculations of N₂O₅: a review. *Aerosol Sc. Technol.*, **45**, 655 - 685,
575 <https://doi.org/10.1080/02786826.2010.551672>, 2011.

576 Defino, R. J., Siotuas, C., and Malik, S.: Potential role of ultrafine particles in
577 associations between airborne particle mass and cardiovascular health. *Environ.*
578 *Health Perspect.*, **113**, 934-938, <https://doi.org/10.1289/ehp.7938>, 2005.

579 Fountoukis, C., and Nenes, A.: ISORROPIA II: a computationally efficient
580 thermodynamic equilibrium model for K⁺-Ca²⁺-Mg²⁺-NH₄⁺-Na⁺-SO₄²⁻-NO₃⁻-Cl⁻-
581 H₂O. *Atmos. Chem. Phys.*, **7**, 4639-4659, [https://doi.org/10.5194/acp-7-4639-](https://doi.org/10.5194/acp-7-4639-2007)
582 2007, 2007.

583 Ge, X., Li, L., Chen, Y., Chen, H., Wu, D., Wang, J., Xie, X., Ge, S., Ye, Z., Xu, J.,
584 and Chen, M.: Aerosol characteristics and sources in Yangzhou, China resolved
585 by offline aerosol mass spectrometry and other techniques. *Environ. Pollut.*, **225**,
586 74-85, <https://doi.org/10.1016/j.environpol.2017.03.044>, 2017.

587 Goodman, A. L., Underwood, G. M., and Grassian, V. H.: A laboratory study of the
588 heterogeneous reaction of nitric acid on calcium carbonate particles. *J. Geophys.*
589 *Res. Atmos.*, **105**, 29053-29064, <https://doi.org/10.1029/2000JD900396>, 2000.

590 Griffith, S. M., Huang, X. H. H., Louie, P. K. K., and Yu, J. Z.: Characterizing the
591 thermodynamic and chemical composition factors controlling PM_{2.5} nitrate:
592 Insights from two years of online measurements in Hong Kong. *Atmos. Environ.*,
593 **122**, 864-875, <https://doi.org/10.1016/j.atmosenv.2015.02.009>, 2015.

594 Guo, H., Xu, L., Bougiatioti, A., Cerully, K. M., Capps, S. L., He, J. R., Carlton,
595 A. G., Lee, S.-H., Bergin, M. H., Ng, N. L., Nenes, A., and Weber, R. J.: Fine
596 particle water and pH in the southeastern United States. *Atmos. Chem. Phys.*, **15**,

597 5221-5228. <https://doi.org/10.5194/acp-15-5211-2015>, 2015.

598 Guo, H., Otjes, R., Schlag, P., Kiendler-Schar, A., Nenes, A., and Weber, R. J.:
599 Effectiveness of ammonia reduction of control of fine particulate nitrate. *Atmos.*
600 *Chem. Phys.*, **18**, 12241-12256. <https://doi.org/10.5194/acp-18-12241>, 2018,
601 2018.

602 He, K., Zhao, Q., Ma, Y., Duan, F., Yang, F., Shi, Z., and Chen, G.: Spatial and
603 seasonal variability of PM_{2.5} acidity at two Chinese megacities: insights into the
604 formation of secondary inorganic aerosols. *Atmos. Chem. Phys.*, **12**, 1377-1395.
605 <https://doi.org/10.5194/acp-12-1377-2012>, 2012.

606 Huang, R.-J., Zhang, Y., Bozzetti, C., Ho, K.-F., Cao, J.-J., Han, Y., Daellenbach, R.,
607 Slowik, J. G., Platt, S. M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Bruns,
608 E. A., Crippa, M., Ciarelli, G., Piazzalunga, A., Schwikkowski, M., Abbaszade,
609 G., Schnelle-Kreis, J., Zimmerman, R., An, Z., Szidat, S., Baltensperger, U.,
610 Haddad, I. E., and Prévôt, A. H.: High secondary aerosol contribution to
611 particulate pollution during haze events in China. *Nature*, **514**, 218-222,
612 <https://doi.org/10.1038/nature13774>, 2014a.

613 Huang, R. J., Chen, R., Jing, M., Yang, L., Li, Y., Chen, Q., Chen, Y., Yan, J., Lin,
614 C., Wu, Y., Zhang, R., Haddad, J. E., Prevot, A. S. H., O'Dowd, C. D., and Cao,
615 J.: Source-specific health risk analysis on particulate trace elements: coal
616 combustion and traffic emission as major contributors in wintertime Beijing.
617 *Environ. Sci. Technol.* **52**, 10967-10974, <https://doi.org/10.1021/acs.est.8b02091>.

618 Huang, Y., Shen, H., Chen, H., Wang, R., Zhang, Y., Su, S., Chen, Y., Lin, N., Zhong,
619 Q., Wang, X., Liu, J., Li, B., Liu, W., and Tao, S.: Quantification of global primary
620 emissions of PM_{2.5}, PM₁₀ and TSP from combustion and industrial process sources.
621 *Environ. Sci. Technol.*, **48**, 13834-13843, <https://doi.org/10.1021/es503696k>,
622 2014b.

623 Hui, L., Liu, X., Tan, Q., Feng, M., An, J., Qu, Y., Zhang, Y., and Jiang, M.:
624 Characteristics, source apportionment and contribution of VOCs to ozone
625 formation in Wuhan, Central China. *Atmos. Environ.*, **192**, 55-71,
626 <https://doi.org/10.1016/j.atmosenv.2018.08.0642>, 2018.

627 Khoder, M. I.: Atmospheric conversion of sulfur dioxide to particulate sulfate and
628 nitrogen dioxide to particulate nitrate and gaseous nitric acid in an urban area.
629 *Chemosphere*, **49**, 675-684, [https://doi.org/10.1016/S0045-6535\(02\)00391-0](https://doi.org/10.1016/S0045-6535(02)00391-0),
630 2002.

631 Kong, L., Yang, Y., Zhang, S., Zhao, X., Du, H., Fu, H., Zhang, S., Cheng, T., Yang,
632 X., Chen, J., Wu, D., Sheng, J., Hong, S., and Jiao, L.: Observation of linear
633 dependence between sulfate and nitrate in atmospheric particles. *J. Geophys. Res.*
634 *Atmos.*, **119**, 341-361, <https://doi.org/10.1002/2013JD020222>, 2014.

635 Lin, Y.-C., and Cheng, M.-T.: Evaluation of formation rates of NO₂ to gaseous and
636 particulate nitrate in the urban atmosphere. *Atmos. Environ.*, **41**, 1903-1910,
637 <https://doi.org/10.1016/j.atmosenv.2006.10.065>, 2007.

638 Lin, Y.-C., Cheng, M.-T., Ting, W.-Y., and Yeh, C.-R.: Characteristics of gaseous
639 HNO₂, HNO₃, NH₃ and particulate ammonium nitrate in an urban city of central
640 Taiwan. *Atmos. Environ.*, **40**(25), 4725-4733,
641 <https://doi.org/10.1016/j.atmosenv.2006.04.037>, 2006.

642 Liu, M., Song, Y., Zhou, T., Xu, Z., Yan, C., Zheng, M., Wu, Z., Hu, M., Wu, Y., and
643 Zhu, T.: Fine particle pH during severe haze episodes in northern China.
644 *Geophys. Res. Lett.*, **44**, 5213-5222, <https://doi.org/10.1002/2017GL073210>,
645 2017.

646 Mental, T. F., Sohn, M., and Wahner, A.: Nitrate effect in the heterogeneous
647 hydrolysis of dinitrogen pentoxide on aqueous aerosols. *Phys. Chem. Chem.*
648 *Phys.*, **1**, 5451-5457, <https://doi.org/10.1039/a905338g>, 1999.

649 Nel, A.: Air pollution-related illness: effects of particles. *Science*, **308**, 804-806,
650 <https://doi.org/10.1126/science.1108752>, 2005.

651 Pan, Y., Tian, S., Zhao, Y., Zhang, L., Zhu, X., Gao, J., Huang, W., Zhou, Y., Song, Y.,
652 Zhang, Q., and Wang, Y.: Identifying ammonia hotspots in China using a national
653 observation work. *Environ. Sci. and Technol.*, **52**, 3926-3934.
654 <https://doi.org/10.1021.acs/est.7b05235>, 2018.

655 Pathak, R. K., Wu, W. S., and Wang, T.: Summertime PM_{2.5} ionic species in four
656 major cities of China: nitrate formation in an ammonia-deficient atmosphere.
657 *Atmos. Chem. and Phys.*, **9**, 1711-1722, <https://doi.org/10.5194/acp-9-1711-2009>,
658 2009.

659 Pathak, R. K., Wang, T., and Wu, W. H.: Nighttime enhancement of PM_{2.5} in
660 ammonia-poor atmospheric conditions in Beijing and Shanghai: Plausible
661 contributions of heterogeneous hydrolysis of N₂O₅ and HNO₃ partitioning.
662 *Atmos. Environ.*, **45**, 1183-1191, <https://doi.org/10.1016/j.atmosenv.2010.09.003>,
663 2011.

664 Polissar, A. V., Hopke, P. K., Paatero, P., Kaufmann, Y. J., Hall, D. K., Bodhaine, B.
665 A., Dutton, E. G., and Harris, J. M.: The aerosol at Barrow, Alaska: long-term
666 trends and source locations. *Atmos. Environ.*, **33**, 2441-2458,
667 <https://doi.org/10.1016/j.atmosenv.2018.08.0642>, 2018.

668 Prabhakar, G., Parworth, C. L., Zhang, X., Kim, H., Young, D. E., Beyersdorf, A. J.,
669 Ziemba, L. D., Nowak, J. B., Bertram, T. H., Faloona, I. C., Zhang, Q., and
670 Cappa, C. D.: Observational assessment of the role of nocturnal residual-layer
671 chemistry in determining daytime surface particulate nitrate concentrations.
672 *Atmos. Chem. Phys.*, **17**, 14747-14770, [https://doi.org/10.5194/acp-17-14747-](https://doi.org/10.5194/acp-17-14747-2017)
673 2017, 2017.

674 Squizzato, S., Masiol, M., Brunelli, A., Pistollato, S., Tarabotti, Z., Rampazzo, G., and

675 Pavoni, B.: Factors determining the formation of secondary inorganic aerosol: a
676 case study in the Po Valley (Italy). *Atmos. Chem. and Phys.*, **13**, 1927-1339,
677 <https://doi.org/10.5194/acp-13-1927-2013>, 2013.

678 Tang, X., Zhang, X., Ci, Z., Guo, J., and Wang, J.: Speciation of the major inorganic
679 salts in atmospheric aerosols of Beijing: China: measurements and comparison
680 with model. *Atmos. Environ.*, **133**, 123-134, [https://doi.org/10.1016/j.atmosenv.](https://doi.org/10.1016/j.atmosenv.2016.03.013)
681 [2016.03.013](https://doi.org/10.1016/j.atmosenv.2016.03.013).

682 Tao, Y., Ye, X., Ma, Z., Xie, Y., Wang, R., Chen, J., Yang, X., and Jiang, S.: Insights
683 into different nitrate formation mechanisms from seasonal variations of secondary
684 inorganic aerosols in Shanghai. *Atmos. Environ.*, **145**, 1-9,
685 <https://doi.org/10.1016/j.atmosenv.2016.09.012>, 2016.

686 ten Brink, H., Otjes, R., Jongejan, P., and Slanina, S.: An instrument for semi-
687 continuous monitoring of the size-distribution of nitrate, ammonium, sulfate and
688 chloride in aerosols. *Atmos. Environ.*, **41**, 2768-2779,
689 <https://doi.org/10.1016/j.atmosenv.2006.11.041>, 2007.

690 Wahner, A., Mental, T. F., Sohn, M., and Stier, J.: Heterogenous reaction of N₂O₅ on
691 sodium nitrate aerosol. *J. Geophys. Res. Atmos.*, **103**, 31103-31112,
692 <https://doi.org/10.1029/1998JD100022>, 1998.

693 Wang, G., Wang, H., Yu, Y., Gao, S., Feng, J., Gao, S., & Wang, L.: Chemical
694 characterization of water-soluble components of PM₁₀ and PM_{2.5} atmospheric
695 aerosols in five locations of Nanjing, China. *Atmos. Environ.*, **37**., 2893-2902.
696 [https://doi.org/10.1016/j.atmosenv.S1352-2310\(03\)00271-1](https://doi.org/10.1016/j.atmosenv.S1352-2310(03)00271-1), 2003.

697 Wang, G., Zhang, R., Geomez, M. E., Yang, L., Zamora, M. L., Hu, M., Lin, Y., Peng,
698 J., Guo, S., Meng, J., Li, J., Cheng, C., Hu, T., Ren, Y., Wang, Y., Gao, J., An, Z.,
699 Zhou, W., Li, G., Wang, J., Tian, P., Marrero-Ortiz, W., Secret, J., Du, Z., Zheng,
700 J., Shang, D., Zheng, L., Shao, M., Wang, W., Huang, Y., Wang, Y., Zhu, Y., Li,

701 Y., Hu, J., Pan, B., Cai, L., Cheng, Y., Ji, Y., Zhang, F., Rosenfeld, D., Liss, P. S.,
702 Duce, R. A., Kolb, C. E., and Molina, M. J.: Persistent sulfate formation from
703 London fog to Chinese haze. *Proc. Natl. Acad. Sci.*, **113**, 13630-13635,
704 [https://doi.org/ 10.1073/pnas.1616540113](https://doi.org/10.1073/pnas.1616540113), 2016.

705 Wang, H., Lu, K., Chen, X., Zhu, Q., Chen, Q., Guo, S., Jiang, M., Li, X., Shang, D.,
706 Tang, Z., Wu, Y., Wu, Z., Zou, Q., Zheng, Y., Zheng, L., Zhu, T., Hu, M., and
707 Zhang, Y.: High N₂O₅ concentrations observed in urban Beijing: implications of a
708 large nitrate formation. *Environ. Sci. Technol. Lett.*, **4**, 416-420,
709 <https://doi.org/10.1021/acsestlett.7b00341>, 2017.

710 Wang, H., Zhu, B., Shen, L., Xu, H., An, J., Xue, G., and Cao, J.: Water soluble ions
711 in atmospheric aerosols measured in five sites in the Yantze River Delta, China:
712 size-fractionated seasonal variation and sources. *Atmos. Environ.*, **123(B)**, 370-
713 379, <https://doi.org/10.1016/j.atmosenv.2015.05.070>, 2015.

714 Wang, W., Yu, J., Cui, Y., He, J., Xue, P., Cao, W., Ying, H., Gao, W., Ying, Y., Gao,
715 W., Yan, Y., Hu, B., Xin, J., Wang, L., Liu, Z., Sun, Y., Ji, D., and Wang, Y.:
716 Characteristics of fine particulate matter and its sources in an industrialized
717 coastal city, Ningbo, Yantze River Delta, China. *Atmos. Res.*, **203**, 105-117,
718 <https://doi.org/10.1016/j.atmosres.2017.11.033>, 2018.

719 Wang, Y., Zhuang, G., Zhang, X., Xu, C., Tang, A., Chen, J., and An, Z.: The ion
720 chemistry, seasonal cycle, and sources of PM_{2.5} and TSP aerosol in Shanghai.
721 *Atmos. Environ.*, **40(16)**, 2935-2952,
722 <https://doi.org/10.1016/j.atmosenv.2005.12.051>, 2006.

723 Wei, L., Yue, S., Zhao, W., Yang, W., Zhang, Y., Ren, L., Han, X., Guo, Q., Sun, Y.,
724 Wang, Z., and Fu, P.: Stable sulfur isotope ratios and chemical compositions of
725 fine aerosols (PM_{2.5}) in Beijing, China. *Sci. Total Environ.*, **633**, 1156-1164,
726 <https://doi.org/10.1016/j.scitotenv.2018.03.153>, 2018.

727 Wen, L., Chen, J., Yang, L., Wang, X., Xu, C., Sui, X., Yao, L., Zhu, Y., Zhang, J.,
728 Zhu, T., and Wang, W.: Enhanced formation of particulate nitrate at a rural site on
729 the North China Plain in summer: the importance roles of ammonia and ozone.
730 *Atmos. Environ.*, **101**, 294-302, <https://doi.org/10.1016/j.atmosenv.2014.11.037>,
731 2015.

732 Yang, H., Yu, J. Z., Ho, S. S. H., Xu, J., Wu, W.-S., Wan, C. H., Wang, X., Wang, X.,
733 and Wang, L.: The chemical composition of inorganic and carbonaceous
734 materials in PM_{2.5} in Nanjing, China. *Atmos. Environ.* , **39**, 3735-3749,
735 <https://doi.org/10.1016/j.atmosenv.2005.03.010>, 2005.

736 Yao, X., Chan, C. K., Fang, M., Cadle, S., Chan, T., Mulawa, P., He, K., and Ye, B.:
737 The water-soluble ionic composition PM_{2.5} in Shanghai and Beijing, China.
738 *Atmos. Environ.*, **36**, 4223-4234, <https://doi.org/10.1016/j.atmosenv.2005.12.051>,
739 2002.

740 Ye, X. N., Ma, Z., Zhang, J. C., Du, H. H., Chen, J. M., Chen, H., Yang, X., Gao, W.,
741 and Geng, F. H.: Important role of ammonia on haze formation in Shanghai.
742 *Environ. Res. Lett.*, **6**, 024019, <https://doi.org/10.1088/1748-9326/6/2024019>,
743 2011.

744 Ye, Z., Liu, J., Gu, A., Feng, F., Liu, Y., Bi, C., Xu, J., Li, L., Chen, H., Chen, Y., Dai,
745 L., Zhou, Q., and Ge, X.: Chemical characterization of fine particulate matter in
746 Changzhou, China and source apportionment with offline aerosol mass
747 spectrometry. *Atmos. Chem. .Phys.*, **17**, 2573-2592, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-17-2573-2017)
748 [17-2573-2017](https://doi.org/10.5194/acp-17-2573-2017), 2017.

749 Zhang, R., Jing, J., Tao, J., Hsu, S.-C., Wang, G., Cao, J., Lee, C.S.L., Zhu, L., Chen,
750 Z., Zhao, Y., and Shen, Z., Chemical characterization and source apportionment
751 of PM_{2.5} in Beijing: seasonal perspective. *Atmos. Chem. .Phys.*, **13**, 7053-7074,
752 <https://doi.org/10.5194/acp-13-7053-2013>, 2013.

753 Zhang, Y.-L., and Cao, F.: Fine particulate matters (PM_{2.5}) in China at a city level. *Sci.*
754 *Rep.*, **5**, 14884, <https://doi.org/10.1038/srep14884>, 2015.

755 Zhao, P. S., Dong, F., He, D., Zhao, X. J., Zhang, X. L., Zhang, W. Z., Yao, Q., and
756 Liu, H. Y.: Characteristics of concentrations and chemical compositions for
757 PM_{2.5} in the region of Beijing, Tianjin, and Hebei, China. *Atmos. Chem. Phys.*,
758 **13**, 4631-4644, <https://doi.org/10.5194/acp-13-4631-2013>, 2013.

759 Zheng, B., Tong, D., Li, M., Liu, F., Hong, C., Geng, G., Li, H., Li, X., Peng, L., Qi,
760 J., Yan, L., Zhang, Y., Zhao, H., Zheng, Y., He, K., and Zhang, Q.: Trends in
761 China's anthropogenic emissions since 2010 as the consequence of clean air
762 actions. *Atmos. Chem. Phys.*, **18**(19), 14095-14111, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-18-14095-2018)
763 [18-14095-2018](https://doi.org/10.5194/acp-18-14095-2018), 2018.

764 Zou, J., Liu, Z., Hu, B., Huang, X., Wen, T., Ji, D., Liu, J., Yang, Y., Yao, and Wang,
765 Y.: Aerosol chemical compositions in the Northern China Plain and the impact on
766 visibility in Beijing and Tianjin., *Atmos. Res.*, **201**, 235-246,
767 <https://doi.org/10.1016/j.atmosres.2017.09.014>, 2018.

768

769 **Table Captions**

770 Table 1 The regression models between $\text{NO}_3^-/\text{SO}_4^{2-}$ (Y) and $\text{NH}_4^+/\text{SO}_4^{2-}$ (X) along
771 with the criterion values of $\text{NH}_4^+/\text{SO}_4^{2-}$ in ammonium-rich regime during the
772 sampling periods.

773

774 **Figure Captions**

775 Figure 1 Time series of concentrations in (a) $\text{PM}_{2.5}$ mass, (b) SIA species, (c) ALWC
776 and (d) Ox along with (e) Fn observed in Nanjing during the sampling
777 periods. The grey shadows represent the high $\text{PM}_{2.5}$ periods discussed in the
778 section 3.6.

779 Figure 2 Abundance of each species in TWSIIs during the (a) entire, (b) haze ($\text{PM}_{2.5} \geq$
780 $150 \mu\text{g m}^{-3}$) and (c) clear ($\text{PM}_{2.5} < 35 \mu\text{g m}^{-3}$) events. The numbers in the
781 parentheses are standard deviations.

782 Figure 3 Abundance of each species in TWSIIs during the (a) entire, (b) haze ($\text{PM}_{2.5} \geq$
783 $150 \mu\text{g m}^{-3}$) and (c) clear ($\text{PM}_{2.5} < 35 \mu\text{g m}^{-3}$) events. The numbers in the
784 parentheses are standard deviations.

785 Figure 4 The PSCF maps of high nitrate pollution.

786 Figure 5 Scatter plots of molar ratios of $\text{NO}_3^-/\text{SO}_4^{2-}$ against $\text{NH}_4^+/\text{SO}_4^{2-}$ in Nanjing
787 during the different seasons.

788 Figure 6 Scatter plot of NO_3^- vs. excess- NH_4^+ molar concentrations in Nanjing during
789 the different seasons. The results in Beijing, Shanghai, Guangzhou, Lanzhou
790 and Hong Kong are also shown in this figure.

791 Figure 7 Scatter plots of (a) Fn against Ox and (b) Fn against ALWC in daytime and
792 nighttime aerosol samples during the high hourly $\text{PM}_{2.5}$ concentration
793 conditions (hourly $\text{PM}_{2.5} \geq 150 \mu\text{g m}^{-3}$).

794 Figure 8 Time series of concentrations in (a) $\text{PM}_{2.5}$ mass and CO, (b) SIA species

795 (NO₃⁻, SO₄²⁻ and NH₄⁺), (c) ALWC, O_x and NO₂ and (d) RH and T in
796 Nanjing City from March 3 to 6, 2016. The grey shadows denote PM_{2.5}
797 episodes. The red numbers represent NO₃⁻ production rate during the PM_{2.5}
798 episodes.

799 Figure 9 Nitrate concentrations simulated by ISORROPIA II model depending on
800 TN and TA concentrations under (a) SO₄²⁻ = 10 μg m⁻³ and (b) SO₄²⁻ = 60
801 μg m⁻³. The purple dots denote the observed TN and TA concentrations at
802 the receptor site during the sampling periods.

803

Table 1 The regression models between $\text{NO}_3^-/\text{SO}_4^{2-}$ (Y) and $\text{NH}_4^+/\text{SO}_4^{2-}$ (X) along with the criterion values of $\text{NH}_4^+/\text{SO}_4^{2-}$ in ammonium-rich regime during the sampling periods.

Sampling periods	Regression models	Criterion values of $\text{NH}_4^+/\text{SO}_4^{2-}$
2016 spring	$Y = 0.71 X - 1.27; R^2 = 0.87$	1.8
2016 summer	$Y = 0.67 X - 1.22; R^2 = 0.86$	1.8
2017 winter	$Y = 0.81 X - 1.50; R^2 = 0.91$	1.9
2017 spring	$Y = 0.95 X - 1.91; R^2 = 0.94$	2.0
2017 summer	$Y = 0.79 X - 1.32; R^2 = 0.84$	1.7

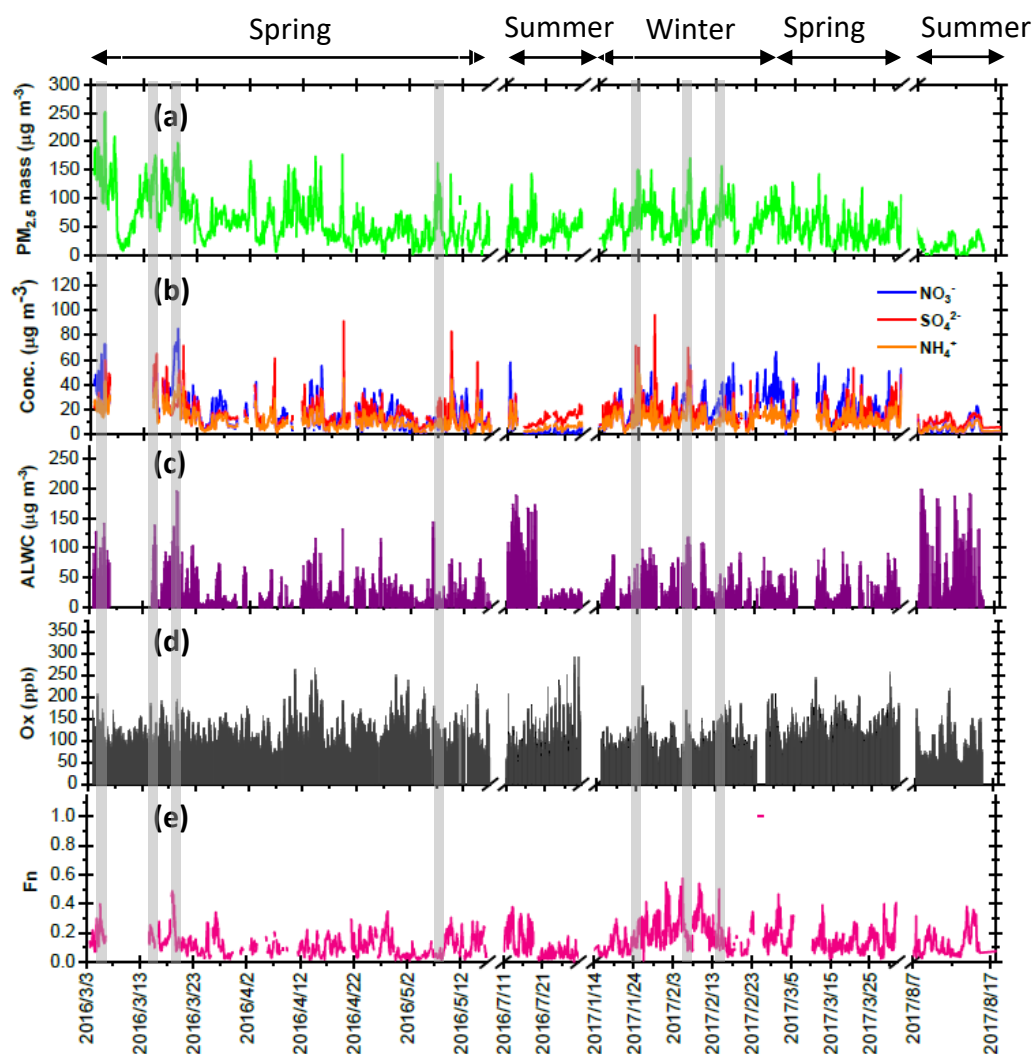
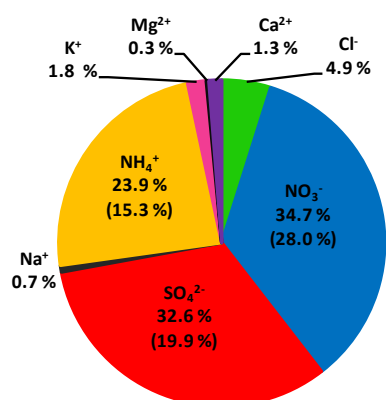
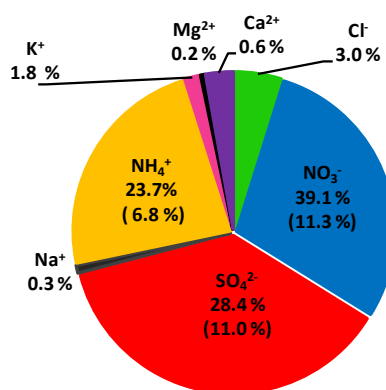


Figure 1 Time series of concentrations in (a) $PM_{2.5}$ mass, (b) SIA species, (c) ALWC and (d) Ox along with (e) Fn observed in Nanjing during the sampling periods. The grey shadows represent the high $PM_{2.5}$ periods discussed in the section 3.6.

(a) Entire days: $PM_{2.5} = 58 \pm 35 \mu g m^{-3}$



(b) Haze events: $PM_{2.5} = 171 \pm 18 \mu g m^{-3}$



(c) Clear events: $PM_{2.5} = 22 \pm 9 \mu g m^{-3}$

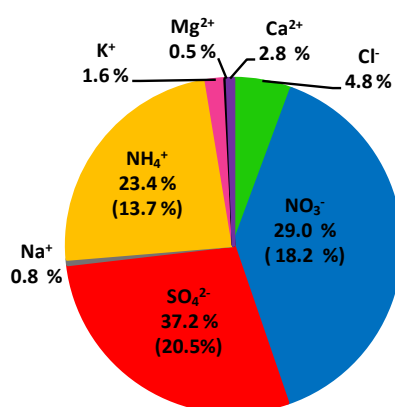


Figure 2 Abundance of each species in TWSIIs during the (a) entire, (b) haze ($PM_{2.5} \geq 150 \mu g m^{-3}$) and (c) clear ($PM_{2.5} < 35 \mu g m^{-3}$) events. The numbers in the parentheses are standard deviations.

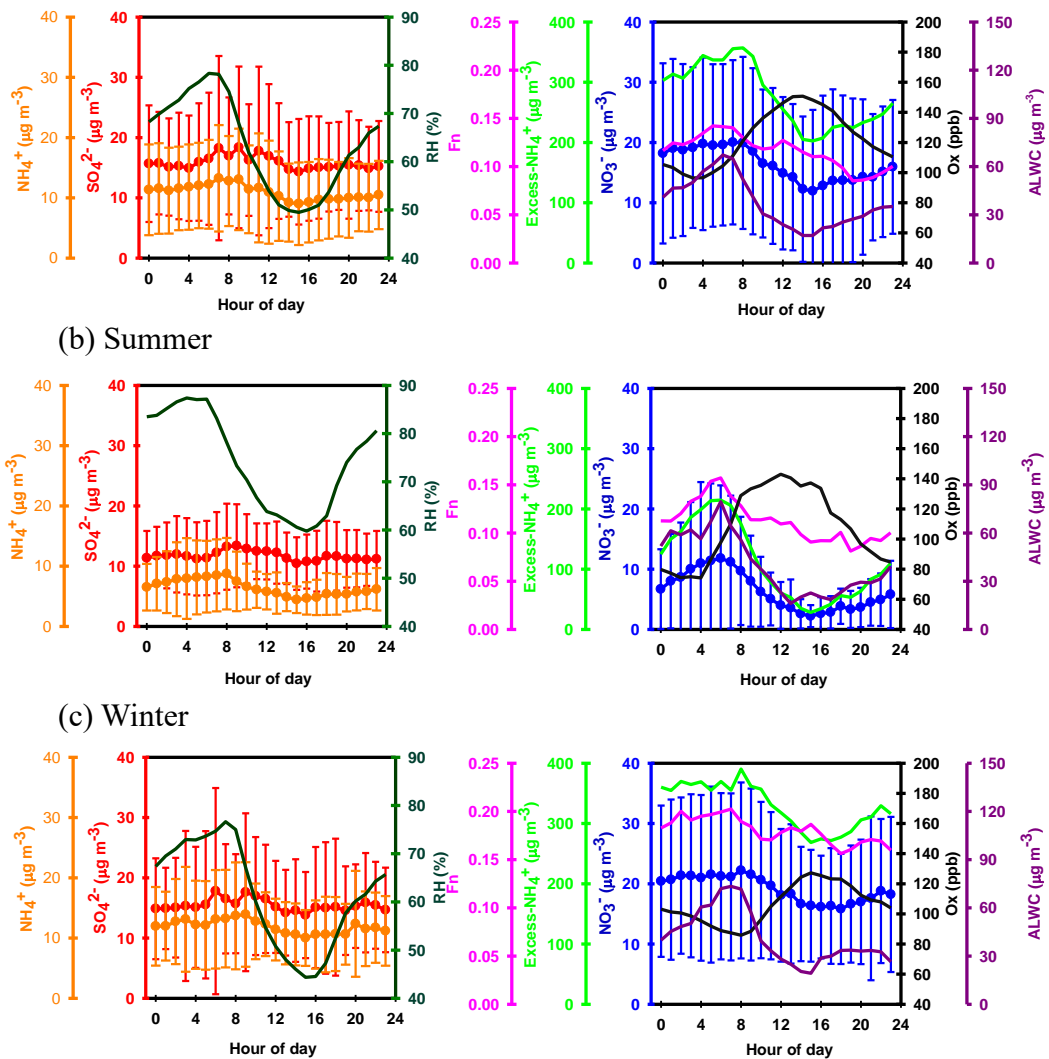


Figure 3 Diurnal variations of the concentrations of NO_3^- , SO_4^{2-} and NH_4^+ , excess- NH_4^+ , Ox and ALWC, and nitrogen conversion ratio (Fn) as well as ambient relative humidity in Nanjing during the sampling periods. For SO_4^{2-} , NO_3^- and NH_4^+ , the mean values (dots) and standard deviations (solid lines) are plotted.

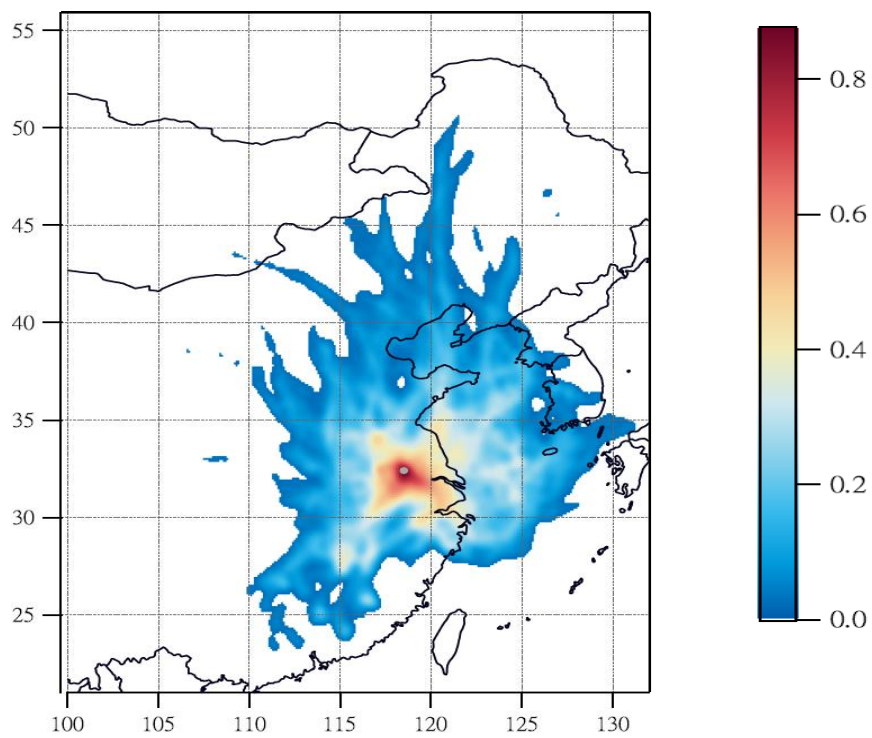


Figure 4 The PSCF maps of high nitrate pollution.

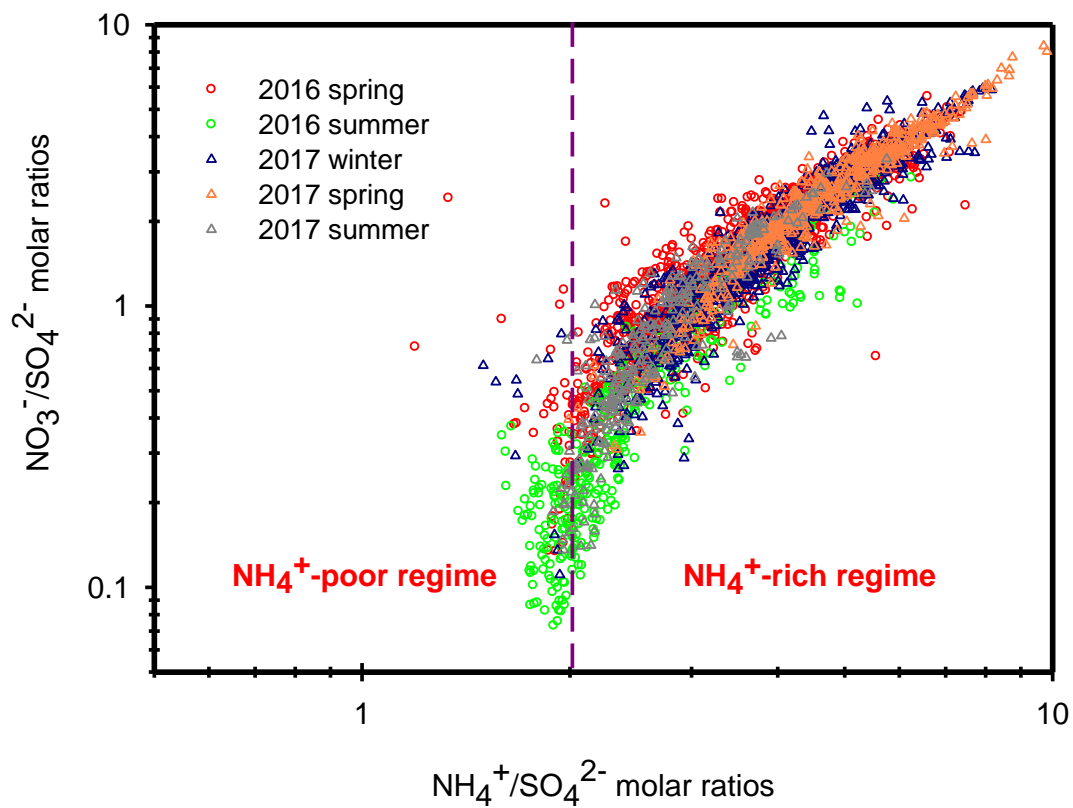


Figure 5 Scatter plots of molar ratios of $\text{NO}_3^-/\text{SO}_4^{2-}$ against $\text{NH}_4^+/\text{SO}_4^{2-}$ in Nanjing during the different seasons.

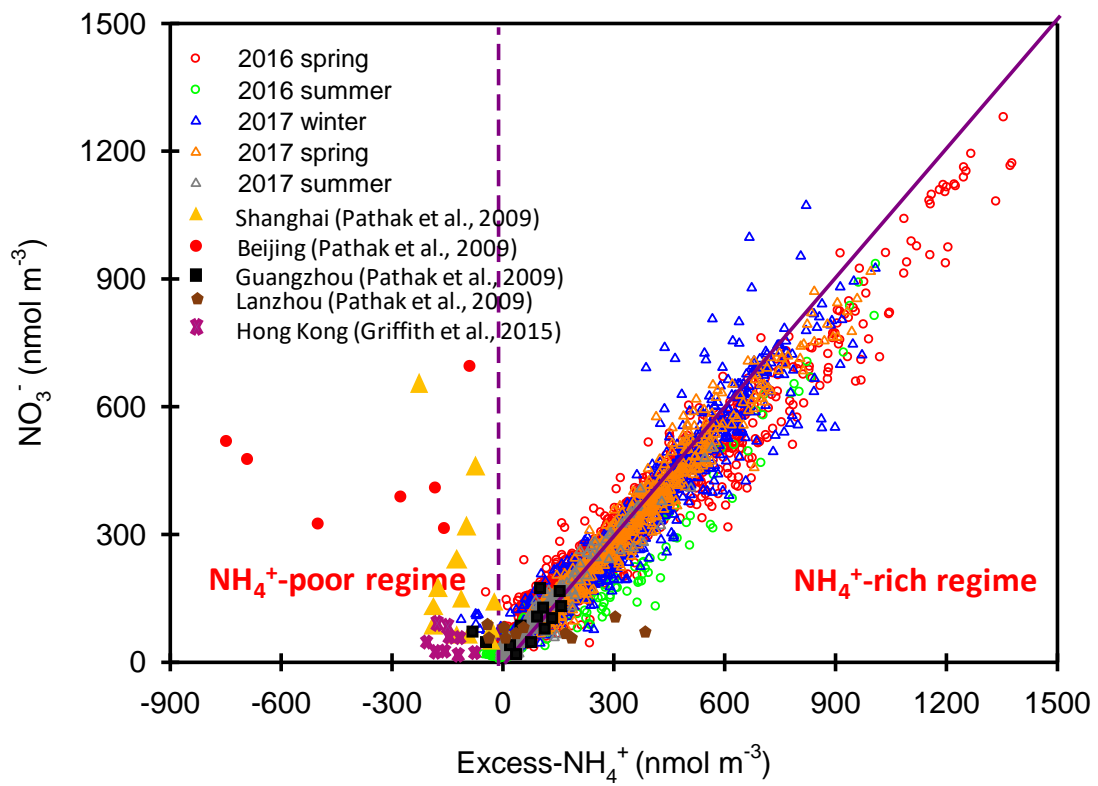


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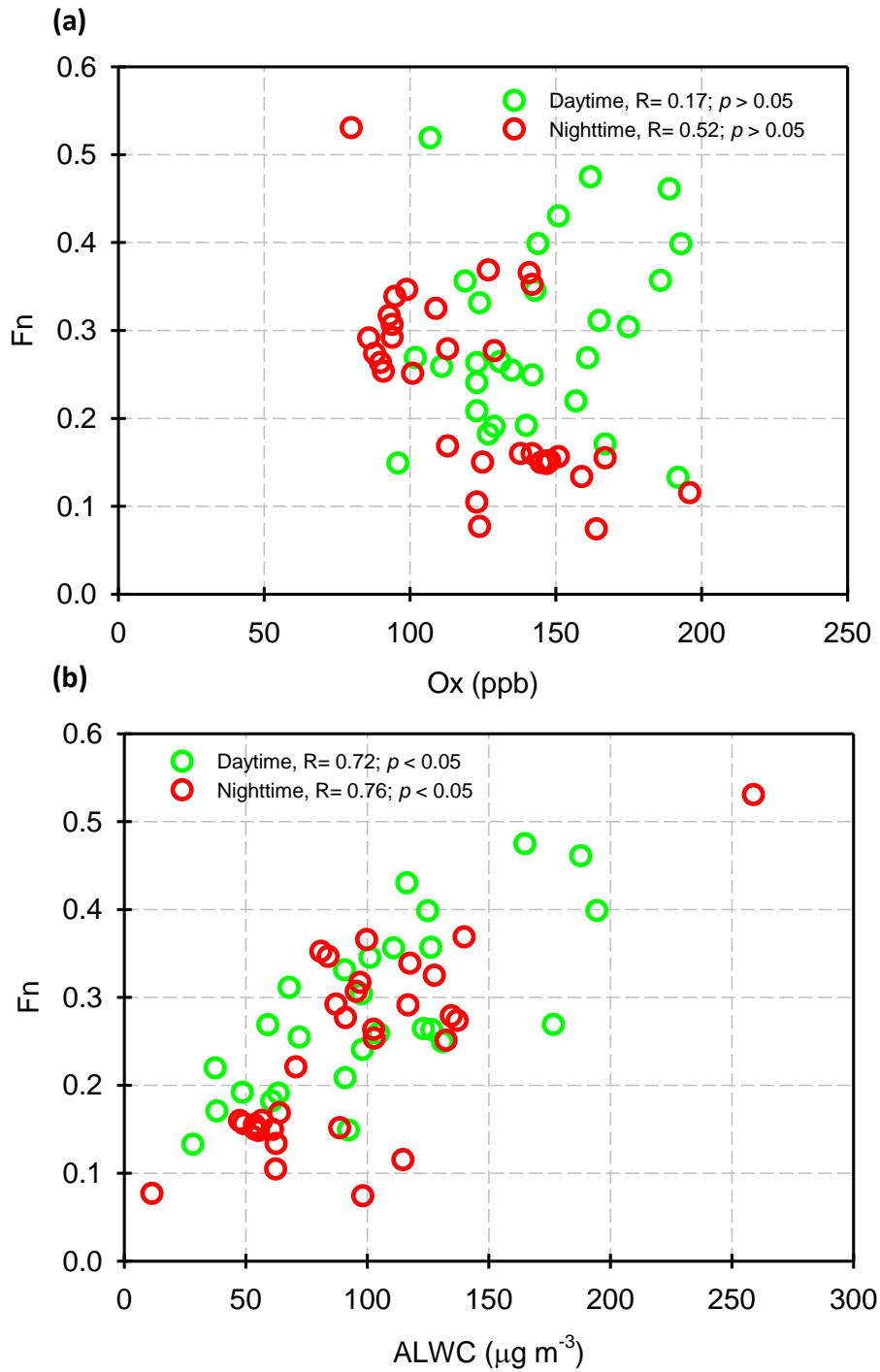


Figure 7 Scatter plots of (a) F_n against Ox and (b) F_n against ALWC in daytime and nighttime aerosol samples during the high hourly PM_{2.5} concentration conditions (hourly PM_{2.5} ≥ 150 µg m⁻³).

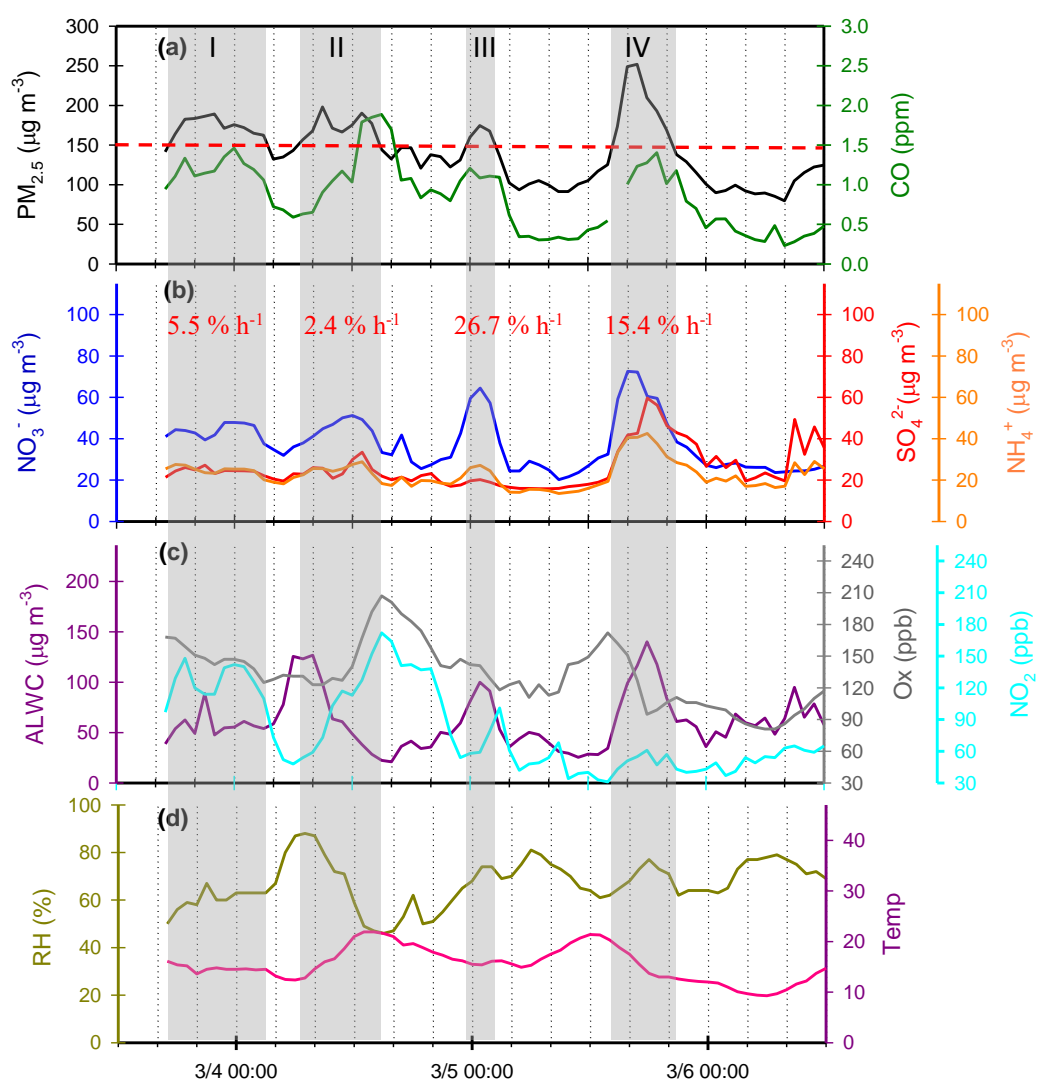


Figure 8 Time series of concentrations in (a) $\text{PM}_{2.5}$ mass and CO, (b) SIA species (NO_3^- , SO_4^{2-} and NH_4^+), (c) ALWC, Ox and NO_2 and (d) RH and T in Nanjing City from March 3 to 6, 2016. The grey shadows denote $\text{PM}_{2.5}$ episodes. The red numbers represent NO_3^- production rate during the $\text{PM}_{2.5}$ episodes.

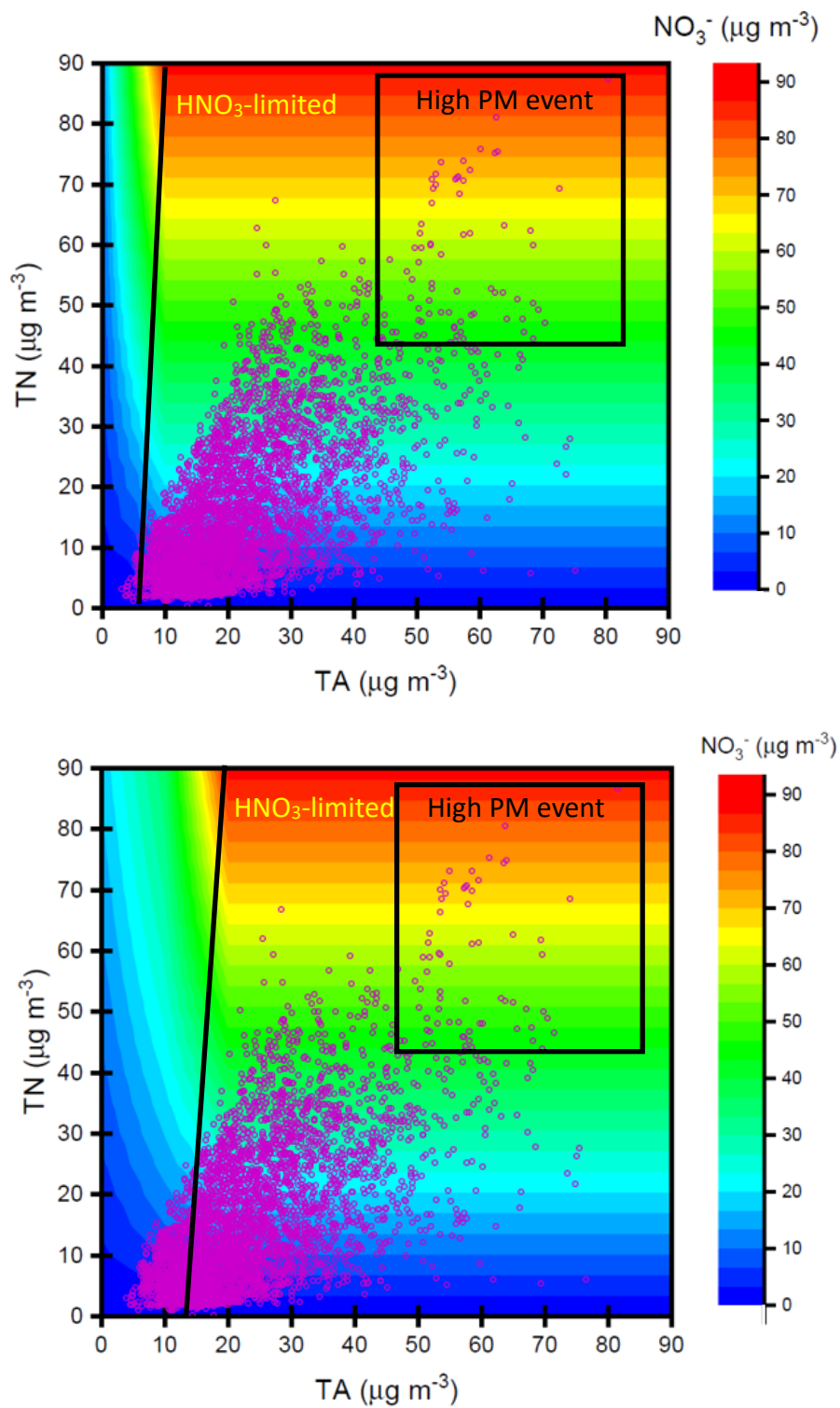


Figure 9 Nitrate concentrations simulated by ISORROPIA II model depending on TN and TA concentrations under (a) $\text{SO}_4^{2-} = 10 \mu\text{g m}^{-3}$ and (b) $\text{SO}_4^{2-} = 60 \mu\text{g m}^{-3}$. The purple dots denote the observed TN and TA concentrations at the receptor site during the sampling periods.