Terpenes and their oxidation products in the French Landes forest: insight from Vocus PTR-TOF measurements

Haiyan Li[1], Matthieu Riva[2], Pekka Rantala[1], Liine Heikkinen[1], Kaspar Daellenbach[1], Jordan E. Krechmer[3], Pierre-Marie Flaud[4,5], Douglas Worsnop[3], Markku Kulmala[1], Eric Villenave[4,5], Emilie Perraudin[4,5], Mikael Ehn[1], Federico Bianchi[1]

[1] Institute for Atmospheric and Earth System Research / Physics, Faculty of Science, University of Helsinki, Finland

[2] Univ. Lyon, Université Claude Bernard Lyon 1, CNRS, IRCELYON, F-69626, Villeurbanne, France

[3] Aerodyne Research Inc., Billerica, Massachusetts 01821, USA

[4] Univ. Bordeaux, EPOC, UMR 5805, F-33405 Talence Cedex, France

[5] CNRS, EPOC, UMR 5805, F-33405 Talence Cedex, France

Correspondence: Haiyan Li (haiyan.li@helsinki.fi) and Matthieu Riva (matthieu.riva@ircelyon.univ-lyon1.fr)

Abstract. The capabilities of the recently developed Vocus proton-transfer-reaction time-of-flight mass spectrometer (PTR-TOF) are reported for the first time based on ambient measurements. With the deployment of the Vocus PTR-TOF, we present an overview of the observed gas-phase (oxygenated) molecules in the French Landes forest during summertime 2018 and gain insights into the atmospheric oxidation of terpenes, which are emitted in large quantities in the atmosphere and play important roles in secondary organic aerosol production. Due to the greatly improved detection efficiency compared to conventional PTR instruments, the Vocus PTR-TOF identifies a large amount of gas-phase signals with elemental composition categories including CH, CHO, CHN, CHS, CHON, CHOS, and others. Multiple hydrocarbons are detected, with carbon numbers up to 20. Particularly, we report the first direct observations of low-volatility diterpenes in the ambient air. The diurnal cycle of diterpenes is similar to that of monoterpenes and sesquiterpenes, but contrary to that of isoprene. Various types of terpene reaction products and intermediates are also characterized. Generally, the more oxidized products from terpene oxidations show a broad peak in the day due to the strong photochemical effects, while the less oxygenated products peak in the early morning and/or in the evening. To evaluate the importance of different formation pathways in terpene chemistry, the reaction rates of terpenes with main oxidants (i.e., hydroxyl radical, OH; ozone, O_3; and nitrate radical, NO_3) are calculated. For the less oxidized non-nitrate monoterpene oxidation products, their morning and evening peaks have contributions from both O_3- and OH-initiated monoterpene oxidation. For the monoterpene-derived organic nitrates, oxidations by O_3, OH, and NO_3 radicals all contribute to their formation, with their relative roles varying considerably over the course of the day. Through a detailed analysis of terpene chemistry, this study demonstrates the capability of the Vocus PTR-TOF in the detection of a wide range of oxidized reaction products in ambient and remote conditions, which highlights its importance in investigating atmospheric oxidation processes.

1. Introduction

Organic aerosol (OA) constitutes a large fraction of atmospheric particles, having significant impacts on climate change, air quality, and human health (Maria et al., 2004; IPCC, 2013; Mauderly and Chow, 2008). On a global scale, secondary OA (SOA) is the largest source of OA, formed through the oxidation of volatile organic compounds (VOCs) (Jimenez et al., 2009). Biogenic VOCs (BVOCs) are released into the atmosphere in high amounts, with an annual global budget being 760 Tg C (Sindelarova et al., 2014). On average, SOA production from biogenic precursors ranges from 2.5 to 44.5 Tg C annually, which is much larger than that from anthropogenic sources (Tsigaridis and Kanakidou, 2003). Over the past decades, a considerable amount of studies has been conducted to investigate the atmospheric chemistry of BVOCs (Kanakidou et al., 2005; Henze et al., 2006; Hatfield et al., 2011; Calfapietra et al., 2013; Jokinen et al., 2015; Ng et al., 2017). However, an incomplete

understanding of BVOCs characteristics and their oxidation processes in the atmosphere remains and yields large uncertainties in quantitative estimates of air quality and climate effects of atmospheric aerosols (Carslaw et al., 2013; Zhu et al., 2019).

Terpenes make up the main fraction of BVOCs (Guenther et al., 1995), encompassing isoprene (C_5H_8), monoterpenes ($C_{10}H_{16}$), sesquiterpenes ($C_{15}H_{24}$), diterpenes ($C_{20}H_{32}$) and even larger compounds. With one or more C=C double bonds in their molecular structures, terpenes are highly reactive. After entering the atmosphere, terpenes can undergo oxidative chemistry with the common atmospheric oxidants including hydroxyl radical (OH), ozone (O_3), and nitrate radical (NO_3). These oxidation processes generate a large variety of organic species, with volatilities ranging from gas-phase volatile species (VOC), to semi-volatile / low volatility organic compounds (SVOC and LVOC), to extremely low volatility organic compounds (ELVOC) and even ultra-low volatility organic compounds (ULVOC), which irreversibly contribute to SOA formation (Donahue et al., 2012). Due to the chemical complexity and low concentrations of BVOCs oxidation products, it remains extremely challenging to provide a comprehensive understanding of terpene chemistry in the atmosphere.

With a high time response and sensitivity, proton-transfer-reaction mass spectrometry (PTR-MS) has been widely used to study the emissions and chemical evolution of VOCs in the atmosphere (Yuan et al., 2017). However, due to the relatively low sensitivity, previous PTR-MS instruments were not optimized to detect low volatility compounds. For example, only a few ambient PTR-MS observations of sesquiterpenes are available (Kim et al., 2009; Jardine et al., 2011). Correspondingly, it is not surprising that ambient observations of diterpenes, which are generally considered to be non-volatile compounds, have never been reported. In addition, the existing PTR-MS is often not sensitive enough to quantify terpene oxidation products at atmospherically relevant concentrations (Yuan et al., 2017). To address these instrumental limitations, two new versions of PTR were recently developed, the PTR3 (Breitenlechner et al., 2017) and the Vocus PTR-TOF (Krechmer et al., 2018), both coupled with a time of flight (TOF) mass analyzer. With the enhanced sensitivities by a factor of ~10 (Holzinger et al., 2019), these instruments are capable of detecting broader spectrum of VOCs, where the detection of low-volatility VOCs is significantly improved compared to the conventional PTR-MS. Based on the laboratory evaluation by Riva et al. (2019a), the Vocus PTR-TOF is able to measure both monoterpenes and lots of monoterpene oxidation products containing up to 6 oxygen atoms.

Known for strong monoterpene emitters (Simon et al., 1994), the Landes forest in southwestern France is a suitable place to investigate atmospheric terpene chemistry. A previous study at this site reported a high nocturnal monoterpene loading and suggested that monoterpene oxidations play an important role in formation of new particles and the consequent growth of atmospheric particles (Kammer et al., 2018). To better assess the roles of BVOCs in aerosol formation, the Characterization of Emissions and Reactivity of Volatile Organic Compounds in the Landes Forest (CERVOLAND campaign) took place in July 2018. The recently developed Vocus PTR-TOF was deployed in the CERVOLAND campaign to characterize terpenes and their gas-phase oxidation products, which provides the first Vocus PTR-TOF measurement in a forested environment. In this work, we present a comprehensive summary of the identified gas-phase molecules and gain insights into terpene chemistry to demonstrate the Vocus PTR-TOF capabilities and the importance of its applications in atmospheric sciences. Characterizations of isoprene, monoterpenes, sesquiterpenes, and particularly the rarely detected diterpenes, are reported. By comparing the reaction rates of different formation pathways, we explore the formation mechanisms of terpene oxidation products, including both non-nitrate and organic nitrate compounds.

2. Experimental methods

2.1 Measurement site

The Vocus PTR-TOF measurements were performed from 8 to 20 July, 2018 in the Landes forest (44º29'39.69"N, 0º57'21.75"W), as part of the CERVOLAND field campaign. The sampling site is situated at the European Integrated Carbon Observation System (ICOS) station at Bilos in southwestern France along the Atlantic coast, ~40 km southwest from the

nearest urban area of the Bordeaux metropole. Both population density and industrial emissions are low in this area. Due to the proximity of the Atlantic Ocean, the site has a strong maritime influence. The forest is largely composed of maritime pines (Pinus pinaster Aiton) and has an average height of ~10m. Monoterpenes are known to be strongly emitted in the forest (Simon et al., 1994), which provides a good place for BVOCs characterization. A more detailed description of the site has been provided in earlier studies (Moreaux et al., 2011; Kammer et al., 2018; Bsaibes et al., 2019).

2.2 Instrumentation

Compared to the conventional PTR instrument, the Vocus PTR-TOF used in this study is mainly differentiated in the following aspects:

1. a new chemical ionization source with a low-pressure reagent-ion source and focusing ion-molecule reactor (FIMR),
2. no dependence of the sensitivity on ambient sample humidity due to the high water mixing ratio (10-20 % v/v) in the FIMR,
3. employment of a TOF mass analyzer with a longer flight tube and faster sampling data acquisition card (mass resolving power up to 15 000 m/dm),
4. an enhanced inlet and source design that minimizes contact between analyte molecules and inlet/source walls, enabling detection of semi- and low-volatility compounds in a similar manner as chemical ionization mass spectrometer (CIMS) instruments (Liu et al. 2019).

Details about the Vocus PTR-TOF are well described by Krechmer et al. (2018). Compared to the ionization in a conventional PTR-MS at 2.0-4.0 mbar, a nitrate CIMS at ambient pressure, and an iodide CIMS at around 100 mbar, the Vocus ionization source is generally operated at a low pressure (Krechmer et al., 2018). In this work, we operated the Vocus ionization source at a pressure of 1.5 mbar. During the campaign, the Vocus PTR-TOF measurements were performed at around 2 m above ground level (a.g.l), thus within the canopy. Sample air was drawn in through 1-m long PTFE tubing (10 mm o.d., 8 mm i.d.) with a flow rate of 4.5 L min^{-1}, which helped to reduce inlet wall losses and sampling delay. Of the total sample flow, only 150 sccm went into the Vocus, while the remainder was directed to the exhaust. The design of FIMR consists of a glass tube with a resistive coating on the inside surface and four quadrupole rods mounted radially on the outside. With an RF field, ions are collimated to the central axis, improving the detection efficiency of product ions. The mass resolving power of the 1.2 m long TOF mass analyzer was 12 000-13 000 m/dm during the whole campaign. Data were recorded with a time resolution of 5 s. Background measurements using high purity nitrogen (UHP N_2) were automatically performed every hour.

The temperature, relative humidity (RH), wind speed, and ambient pressure were continuously monitored at 3.4 m a.g.l. whereas the solar radiation was measured at 15.6 m a.g.l from a mast located at the site. The mixing ratios of nitrogen oxides (NO_x) and ozone (O_3) were measured at 4 m a.g.l with UV absorption and chemiluminescence analyzers, respectively. All data are reported in Coordinated Universal Time (UTC).

2.3 Data analysis and quantification of multiple compounds

Data analysis was performed using the software package “Tofware” (https://www.tofwerk.com/software/tofware/) that runs in the Igor Pro environment (WaveMetrics, OR, USA). Tofware enables the time-dependent mass calibration, baseline subtraction, and assignment of a molecular formula to the identified ions by high resolution analysis. Signals were averaged over 30 min before mass calibration. Due to the high resolving power of the LTOF mass analyzer, isobaric ions were more clearly separated. Examples of peak identification are given in Fig. S1.

The Vocus was calibrated twice a day during the campaign with a mixture (70 ppb each) of terpenes (*m/z* 137: alpha/beta pinene + limonene; *m/z* 135: *p*-cymene) that was diluted using UHP N_2. Similar to conventional PTR instruments, the sensitivities of different VOCs in the Vocus PTR-TOF are linearly related to their proton-transfer reaction rate constants

(*k*) when ion transmission efficiency and fragmentation ions are considered (Sekimoto et al., 2017; Krechmer et al., 2018). Krechmer et al. (2018) have shown that within the Vocus PTR-TOF, the transmission efficiencies of ions > *m/z* 100 Th reach up to 99%. Therefore, the influence of fragmentation correction should be included in this study. According to terpene calibrations, the residual fraction was on average 66% and 55%, respectively, for protonated monoterpenes and *p*-cymene after their fragmentation within the instrument. Based on the corrected sensitivities for fragmentation and the *k* values of monoterpenes and *p*-cymene, an empirical relationship between the sensitivity and *k* was built from the scatterplots using linear regression: Sensitivity (cps ppb^{-1}) = 828.9 × *k* (Fig. S2). Once *k* is available, the sensitivity of a compound can be predicted. It should be noted that the established relationship in this study is not applicable to other conditions or instruments. Some studies found that isoprene may fragment significantly to *m/z* 41(Keck et al., 2008; Schwarz et al., 2009). However, with the ambient data in this work, isoprene seems not to fragment much to $C_3H_5^+$, and they correlate poorly with each other (Fig. S3). Therefore, the fragmentation of isoprene is not considered for its quantification. Sesquiterpenes and some terpene oxidation products were found to fragment to varying degrees (Kim et al., 2009; Kari et al., 2018). Due to the lack of calibrations using other terpenes or terpene oxidation products, their fragmentation patterns within the Vocus PTR-TOF are not known in this work. Therefore, all the other terpenes and terpene oxidation products were quantified without consideration of fragment ions, which should be regarded as the lower limit of their ambient concentrations.

Rate constants for the proton-transfer reactions have only been measured for a subset of compounds. To quantify terpenes and their oxidation products, we used the method proposed by Sekimoto et al. (2017) to calculate the rate constants of different compounds with the polarizability and permanent dipole moment of the molecule. According to Sekimoto et al. (2017), the polarizability and dipole moment of a molecule can be obtained based on the molecular mass, elemental composition, and functionality of the compound. For a class of VOCs with the same number of electronegative atoms, their polarizabilities can be well described using their molecular mass (Sekimoto et al., 2017). For VOCs containing a specific functional group, it is found that their dipole moments are relatively constant based on results in the CRC Handbook (Lide, 2005). Since no isomer information is provided by mass spectrometry alone, it is challenging to figure out the functionality of different compounds. Therefore, the polarizability and dipole moment of the compounds observed in this study were estimated only based on the molecular mass and elemental composition. In this work, based on the physical properties of various compounds in CRC Handbook (Lide, 2005) and the results in Sekimoto et al. (2017), we built the functions between polarizability (α) and molecular mass (M_R) for different groups of VOCs and calculated the average dipole moment (μ) for each group. For example, the polarizabilities of hydrocarbons were approximated as $\alpha = 0.142\ M_R - 0.3$ and the dipole moment was approximated to be zero. For the non-nitrate oxygenated compounds with one oxygen, $\alpha = 0.133\ M_R - 1.2$, and the dipole moment was averaged to be 1.6.

It should be noted that uncertainties are introduced to the calculated sensitivities in the following factors. Firstly, the small difference between the rate coefficients of monoterpenes and *p*-cymene may lead to large uncertainty in the established linear regression function between sensitivity and *k*. Calibrations with more VOC compounds should be performed in future works to cover a larger range of *k* values. Secondly, as mentioned above, the theoretically calculated sensitivities of sesquiterpenes, diterpenes, and terpene oxidation products, may be underestimated to varying extent without the consideration of their fragment ions. Further, some low-volatility compounds may experience wall losses inside the inlet tubing and the instrument and therefore have worse transmissions. The method in this work may overestimate the sensitivities of these low-volatility compounds. In addition to proton transfer reactions, some VOCs can be ionized through ligand switching reactions with water cluster ($(H_2O)_nH_3O^+$) (Tani et al., 2004), thus increasing their sensitivity. However, with the calibration standards used in this study, it is hard to estimate the effect of ligand switching ionization. Lastly, uncertainties come from the estimation of polarizability and dipole moment of a molecule. With the method used in this study, the sensitivity is calculated to be within 50% error when only the elemental composition of a compound is known (Sekimoto et al., 2017).

3. Results and discussion

3.1 Meteorology and trace gases

Figure 1 displays the time variations of meteorological conditions and trace gases during the observation period. The weather was mostly sunny, with solar radiation varying from 400 to 800 W/m^2 during daytime, indicating strong photochemical activity. The ambient temperature and RH varied regularly every day. On average, the temperature was 22.8 ± 5.9 °C (mean ± SE), ranging from 12.1 to 35.0 °C, which is favorable for BVOCs emissions in the forest. The average RH was 70.5 ± 19.0 % during the campaign. Generally, the air masses were quite stable within the canopy. The wind speed never exceeded 1 m/s, indicating the major influence of local sources on atmospheric processes in this study.

The O_3 levels fluctuated dramatically between day and night during the campaign. The average O_3 diurnal cycle showed that O_3 concentration peaked up to ~50 ppb in the daytime. However, during most of the nights, O_3 concentration dropped below 2 ppb. Considering the high nighttime concentration of terpenes observed by the previous study at this site in the same season (Kammer et al., 2018), the low O_3 level at night suggests the large consumption of O_3 by terpenes. Such reactions of terpenes with O_3 can produce low volatility organic compounds, thus contributing to SOA formation (Presto et al., 2005; Jokinen et al., 2014). In addition, plant surface uptake is likely another important ozone sink in the canopy (Goldstein et al., 2004).

The NO concentration was generally low during the campaign, below detection limit (i.e., <0.5 ppb) most of the time. However, clear NO plumes was sometimes observed in the early morning, as shown in Fig.1e. The NO concentration peak at 4 am is probably the combination of local emission sources and low boundary layer. With the increasing sunlight afterwards, the NO concentration started to decrease. A similar diel pattern of NO_2 was observed by the previous study at this site (Kammer et al., 2018). The lower NO_2 concentration during daytime is likely explained by dilution with increasing boundary layer height and NO_2 photolysis.

3.2 Vocus PTR-TOF capabilities in the forest

While Krechmer et al. (2018) and Riva et al. (2019a) have described the novel setup and performance of the Vocus PTR-TOF and its application during a lab study, the instrument capability has not been fully explored in an ambient environment. Based on the CERVOLAND deployment, we provide here, the first overview of gas-phase molecules measured by the Vocus PTR-TOF in the forest. For a better visualization of the complex data set from real atmosphere, mass defect plots (averaged over the whole campaign) are shown in Fig. 2 with the difference between the exact mass and the nominal mass of a compound plotted against its exact mass. With the addition of hydrogen atoms, the mass defect increases, while the addition of oxygen atoms decreases the mass defect. Therefore, changes in the mass defect plot help to provide information on chemical transformation such as oxidation.

The mass defect plot in Fig. 2a is colored according to the retrieved elemental composition, with the black circle indicating unidentified molecules. The size of the markers is proportional to the logarithm of the peak area of the molecule. During the campaign, the Vocus PTR-TOF detected large amounts of (O)VOCs, with elemental composition categories of CH, CHO, CHN, CHS, CHON, CHOS, and others. For hydrocarbons, multiple series with different carbon numbers were measured, especially those compounds containing 5 ("C_5") to 10 carbon atoms ("C_{10}"), 15 carbon atoms ("C_{15}"), and 20 carbon atoms ("C_{20}"). Some of the C5 – C9 ions can be fragments of terpenes and their oxidation products (Tani et al., 2003, 2013; Kim et al., 2009; Kari et al., 2018). For ions <35 Th, the detection efficiency is much reduced due to a high-pass band filter of the BSQ (Krechmer et al., 2018). Compared to the conventional PTR instruments, the observation of larger hydrocarbon molecules by the Vocus PTR-TOF is mainly caused by the much lower wall losses and increased detection efficiency. Hydrocarbon signals were largely contributed by monoterpene ($C_{10}H_{16}H^+$) and its major fragment ($C_6H_8H^+$), indicating the monoterpene-dominated environment in the Landes forest (Kammer et al., 2018). According to previous studies, monoterpene emissions in

the Landes forest are dominated by α-pinene and β-pinene (Simon et al., 1994; Kammer et al., 2018). The identified compound with the elemental composition of $C_4H_9^+$ ranked the third largest peak in hydrocarbons. Detailed discussion about $C_4H_9^+$ ions can be found in the supplement.

In addition to the emitted precursors, the Vocus PTR-TOF detected various VOCs reaction products and intermediates. Similar to the PTR3 measurements in the CLOUD chamber (Breitenlechner et al., 2017), many oxygenated compounds from terpene reactions with varying degrees of oxidation were observed in this study. However, as a potential limitation of the instrument, no dimers in the atmosphere were identified by the Vocus PTR-TOF, consistent with the results from a previous laboratory deployment (Riva et al., 2019a).

Figure 2b compares the daytime and nighttime variations of different molecules, with the marker sized by the signal difference between day and night. The daytime periods cover from 4:30 am to 7:30 pm, and the nighttime periods are from 7:30 pm to 4:30 am of the next day (both are UTC time; Local time = UTC time + 2). The data points are colored in orange when the nighttime signal of the compound is larger than its daytime signal, and in green when the daytime signal is higher. Patterns in the figure clearly show the difference in the diurnal variations of gas molecules with different oxidation degrees. For example, most hydrocarbons are characterized with higher concentrations at night, which is largely caused by the stable nocturnal boundary layer. The more oxidized compounds with more oxygen numbers are generally more abundant during the day due to enhanced photochemistry, whereas the concentrations of the less oxidized compounds are mostly higher at night. Details on the diurnal profiles of different oxidation products and their formation mechanisms are provided in Sect. 3.4.

3.3 Terpene characteristics

The characterizations of isoprene, monoterpenes, sesquiterpenes, and the rarely reported diterpenes, are investigated in this study (Fig. 3, Fig. 4). On the global scale, isoprene is the most emitted BVOC species. It has been well established that photooxidation of isoprene in the atmosphere contributes to SOA formation through the multiphase reactions of isoprene-derived oxidation products (Claeys et al., 2004; Henze and Seinfeld, 2006; Surratt et al., 2010). However, recent advances on isoprene chemistry found that isoprene can impact both particle number and mass of monoterpene-derived SOA by scavenging hydroxyl and peroxy radicals (Kiendler-Scharr et al., 2009; Kanawade et al., 2011; McFiggans et al., 2019). During the CERVOLAND campaign, the average mixing ratio of isoprene was 0.6 ppb, consistent with the mean value of 0.4 ppb reported for the LANDEX campaign during summer 2017 at the same site (Mermet et al., 2019). These values are much lower than that in the southeastern United States (Xiong et al., 2015) and Amazon rainforest (Wei et al., 2018) but higher than observations in the boreal forest at the SMEAR II station (Hellén et al., 2018). Isoprene emissions are strongly light-dependent (Monson et al., 1989; Kaser et al., 2013). Therefore, a pronounced diurnal pattern of isoprene was observed with maximum mixing ratios occurring during daytime and minima at night. It has been shown that the attribution of $C_5H_9^+$ ions to isoprene with PTR instruments can be influenced by the fragmentation of many other compounds, i.e., cycloalkane and 2-methyl-3-buten-2-ol (MBO) (Karl et al., 2012; Gueneron et al., 2015). For example, using an *E/N* ratio of 106 Td in the PTR-MS with a quadrupole mass analyzer, 71% of the parent MBO fragmented to $C_5H_9^+$ ions (Warneke et al., 2003). However, in this study, the $C_5H_9^+$ signal was around 10 times as high as the $C_5H_{11}O^+$ signal and both ions correlated poorly with each other (Fig. S4; $r^2 = 0.33$). This information demonstrate that the fragmentation of MBO does not likely have a significant influence on the attribution of $C_5H_9^+$ ions to isoprene in this work.

As expected, monoterpenes showed the highest mixing ratios among all the terpenes, with an average value of 6.0 ppb. On July 9, a heavy monoterpene episode occurred at night, with the monoterpene mixing ratio reaching as high as 41.2 ppb. Comparatively, the average monoterpene level observed in this work is similar to the measurements performed in 2015 and 2017 at the same site (Kammer et al., 2018; Mermet et al., 2019) and more than ten times higher than that observed in the boreal forest at SMEAR II in summer (Hakola et al., 2012; Hellén et al., 2018). The high concentration of monoterpenes indicates the potential significance of monoterpene-related aerosol chemistry in the Landes forest. Different from the light-

dependence of isoprene emissions, monoterpene emissions are found to be mainly controlled by temperature (Hakola et al., 2006; Kaser et al., 2013). At night, monoterpenes can be continuously emitted and accumulated within the boundary layer. Therefore, monoterpenes showed the opposite diel pattern to isoprene and peaked during nighttime. During daytime, the concentration of monoterpenes dropped to around 0.9 ppb, due to the increased atmospheric mixing after sunrise and the rapid photochemical consumptions.

A study in Hyytiälä concluded that sesquiterpenes, due to their higher reactivity, could play a more important role in O_3 chemistry than monoterpenes, even though the concentration of sesquiterpenes was much lower (Hellén et al., 2018). However, the short lifetimes of sesquiterpenes also mean that their concentrations will be highly dependent on the sampling location at a given site. Some studies also proposed that sesquiterpene oxidation products are linked to atmospheric new particle formation (Bonn and Moortgat, 2003; Boy et al., 2007). Despite the potential importance of sesquiterpenes in aerosol chemistry, the available data on ambient sesquiterpene quantification remains still quite limited. In this work, the mixing ratios of sesquiterpenes were found to vary from 8.9 ppt to 408.9 ppt in the Landes forest, with an average of 64.5 ppt during the observations. This sesquiterpene level is comparable to that reported by Mermet et al. (2019) in summer 2017 at the same site and observations by Jardine et al. (2011) in Amazonia but higher than previous measurements at SMEAR II station (Hellén et al., 2018). Kim et al. (2009) show that different sesquiterpenes fragment on monoterpene parent and fragment ions to varying degrees inside the PTR instruments. Without the consideration of sesuiterpene fragmentation, the quantification of sesquiterpenes in this work may be underestimated. As shown in Fig. 4, sesquiterpenes displayed a similar diurnal pattern with monoterpenes, consistent with observations in other areas (Jardine et al., 2011; Hellén et al., 2018).

While diterpenes are present in all plants in the form of phytol, they have been thought for a long time not to be released by vegetation due to their low volatility (Keeling and Bohlmann, 2006). In 2004, von Schwartzenberg et al. (2004) reported for the first time the release of plant-derived diterpenes into the air. A recent study found that the emission rate of diterpenes by Mediterranean vegetation was in the same order of magnitude as monoterpenes and sesquiterpenes (Yáñez-Serrano et al., 2018). For the first time, this study reports the ambient concentration of diterpenes in a forest. According to the Vocus PTR-TOF measurements, the average mixing ratio of diterpenes was around 2 ppt in the Landes forest. Considering the low volatility of diterpenes and their potential wall losses inside the inlet tubing and the instrument, the diterpene concentration might be higher. Similar to monoterpenes and sesquiterpenes, diterpenes presented peak concentrations at night and lower levels during the day. Although the amounts of diterpenes in the atmosphere are hundreds to thousands times lower than those of monoterpenes and sesquiterpenes, diterpenes potentially play a role in atmospheric chemistry due to their unsaturated structure and high molecular weight (Matsunaga et al., 2012). Up to now, there is no report on the possible atmospheric implications of diterpenes, which should deserve more attention in the future.

Considering the similar atmospheric behaviors of monoterpenes, sesquiterpenes, and diterpenes in this study, it is questioned if the observed sesquiterpenes and diterpenes are real signals in the atmosphere or generated by monoterpenes in the instrument. Bernhammer et al. (2018) have shown that secondary association reactions of protonated isoprene with isoprene can form monoterpenes within the PTR reaction chamber. Figure 5 illustrates the scatter plots among monoterpenes, sesquiterpenes, and diterpenes, colored by time of the day. At night, both sesquiterpenes and diterpenes correlated well with monoterpenes. However, their correlation with monoterpenes got weaker during daytime as the data points became more scattered. This suggests that the observations of sesquiterpenes and diterpenes are real emissions in the atmosphere. Comparatively, sesquiterpenes and diterpenes showed a strong correlation with each other through the whole day ($r^2 = 0.85$).

3.4 Insights into terpene chemistry

3.4.1 Comparison with chamber results

Due to the diverse precursors and changing environmental conditions in the ambient air, it is challenging to retrieve all the atmospheric chemical processes occurring within the Landes forest. To start with, we compare the ambient data with those

from α-pinene ozonolysis in the presence of NO_x conducted in the COALA chamber at the University of Helsinki. A detailed description of the laboratory experiment is provided elsewhere (Riva et al., 2019a, 2019b). According to literature, monoterpenes undergo some degree of fragmentation within the PTR instrument, producing dominant ions of $C_6H_9^+$, $C_5H_7^+$, $C_7H_{11}^+$, et al (Tani et al., 2003, 2013; Kari et al., 2018). As illustrated in Fig. 6, $C_6H_9^+$ is the largest fragment produced by monoterpenes within the Vocus PTR-TOF. However, a clear difference of monoterpene fragmentation pattern is observed in the mass spectra of ambient observations and chamber experiments. While the signal of $C_6H_9^+$ is lower than that of $C_{10}H_{17}^+$ during the field deployment, $C_6H_9^+$ peak is higher than $C_{10}H_{17}^+$ peak in the chamber study. Based on the monoterpene calibration data, the $C_6H_9^+$ signal is around 40% and 138% of the protonated monoterpene signal in ambient deployment and chamber experiment, respectively. The larger presence of the $C_6H_9^+$ peak in the chamber study can be likely explained by the much higher concentrations of oxygenated terpenoids during the chamber experiments. Indeed, previous studies have shown that oxygenated terpenoids, including linalool and pinonaldehyde, fragment inside the PTR instrument and produce a dominant ion at *m/z* 81 (Maleknia et al., 2007; Tani, 2013). Different settings of the instrument can also contribute to different fragmentation patterns of monoterpenes (Tani et al., 2003, 2013; Kari et al., 2018). In our ambient and chamber studies, the *E/N* values of the Vocus PTR-TOF are quite similar, 118 Td and 120 Td, respectively. In addition, the fragmentation patterns vary among individual monoterpene species due to their different physicochemical properties (Tani et al., 2013; Kari et al., 2018). Considering that α-pinene is the only monoterpene species injected in the chamber experiment, the combination of various monoterpenes in the atmosphere likely introduces additional differences in the fragmentation pattern.

Gas-phase ozonolysis of alkenes generates OH radicals in high yields (Rickard et al., 1999). Without an OH scavenger, both O_3- and OH-initiated oxidations happened during α-pinene ozonolysis in the chamber. Using the Vocus PTR-TOF, various oxidation products were identified in the chamber study, with the dominant species being $C_7H_{10,12}O_{3\text{-}6}$, $C_8H_{14}O_{3\text{-}6}$, $C_9H_{14}O_{1\text{-}5}$, and $C_{10}H_{14,16}O_{2\text{-}6}$. In comparison, more oxygenated compounds which were directly emitted or from monoterpene reactions were observed in ambient air due to complex environmental conditions, with the oxygen number ranging from 1 to 7. Therefore, the Vocus PTR-TOF measurements provide the opportunity to characterize both the emitted precursors and the resulting oxidation products. During the chamber experiments, NO_2 was injected and photolyzed using 400nm LED lights to generate NO. In the presence of NO_x, organic nitrates were formed from the reactions between NO and monoterpene-derived peroxy radicals (RO_2). The major organic nitrates observed were $C_9H_{13,15}NO_{6\text{-}8}$ and $C_{10}H_{13,15}NO_{3\text{-}8}$. Compared to the chamber study, more organic nitrates of C_8, C_9, and C_{10} from monoterpene reactions were identified in CERVOLAND data. It is worth pointing out that the combination of different monoterpene species in the ambient environment may result in various types of organic nitrates through different formation pathways.

3.4.2 Non-nitrate terpene oxidation products

Based on the ambient observations, the non-nitrate oxidation products from isoprene, monoterpenes, and sesquiterpenes, are investigated in this study. Isoprene gas-phase products are mainly represented by C_4 and C_5 compounds (Wennberg et al., 2018). In this work, we consider $C_4H_{6,8}O_n$ and $C_5H_{8,10,12}O_n$ (n=1~6) as the dominant non-nitrate products from isoprene oxidations. The diurnal variations of $C_5H_8O_n$ are displayed in Fig. 7 and the others in Fig. S5-7. Generally, all these oxidation products displayed an evening peak at around 8 pm, which may come from the O_3- or OH-initiated isoprene oxidations. Reaction with OH represents the largest loss pathway for isoprene in the atmosphere and produces a population of isoprene peroxyl radicals (Wennberg et al., 2018). In the presence of NO, the major products are methyl vinyl ketone (MVK, C_4H_6O) and methacrolein (MACR, C_4H_6O). Globally, reactions with O_3 contribute a small fraction of approximately 10% to isoprene removal in the atmosphere (Wennberg et al., 2018). When isoprene reacts with O_3, one carbon is always split off from the molecule (Criegee, 1975). Considering the peak concentration of isoprene at 8 pm and the relatively high O_3 concentration at the moment (Figs. 1 and 4), isoprene ozonolysis is also likely contributing to the formation of C_4 oxidation products. Because OH radicals can be efficiently produced from alkene ozonolysis (Pfeiffer et al., 2001), the OH-initiated oxidation of isoprene

can also be an important formation pathway of these oxidation products in the evening. For example, as a predominant product from the reactions of isoprene with OH, $C_5H_{10}O_3$ (corresponding to isoprene hydroxy hydroperoxide and/or isoprene epoxydiols) presented a clear single peak in the evening. To determine the relative importance of O_3- and OH-initiated oxidations in isoprene chemistry at night, the reaction rates (R) of isoprene with O_3 and OH radical were compared by Eq. (1) and Eq. (2):

$$R_{ISO+OH} = k_{ISO+OH}[ISO][OH] \quad (1)$$

$$R_{ISO+O3} = k_{ISO+O3}[ISO][O_3] \quad (2)$$

where k is the reaction rate coefficient of isoprene with OH or O_3, and [ISO], [OH] or [O_3] is the concentration of isoprene, OH radical or O_3.

Taking the evening peak of isoprene oxidation products at 8 pm as an example, we compared the roles of O_3 and OH radicals in their formation. Laboratory studies have shown that the reaction rate coefficient of isoprene with OH radical is generally 10^7 times larger than that of isoprene with O_3 (Dreyfus et al., 2002; Kari et al., 2004). Based on the competition between OH production and removal processes at night (Dusanter et al., 2008), the steady state OH concentration was estimated to be 0.012 ppt. Details can be found in the supplement. With an O_3 concentration of ~20 ppb at 8 pm, the reaction rate of isoprene with OH radical was around 6 times as high as that of isoprene with O_3. For the more oxidized compounds from isoprene oxidations, their concentrations had a broad daytime presence from 10 am to 8 pm due to strong photooxidation processes. Similar diurnal variations of $C_4H_{6,8}O_{5,6}$ and $C_5H_{8,10,12}O_{5,6}$ measured by nitrate CIMS have been observed in an isoprene-dominated environment at Centreville, Alabama (Massoli et al., 2018).

The diurnal patterns of $C_8H_{12,14}O_n$, $C_9H_{14}O_n$, and $C_{10}H_{14,16,18}O_n$ (n=1~6) were illustrated to characterize monoterpene oxidations in the Landes forest (Fig. 8; Fig. S8-12). For the less oxidized compounds with oxygen numbers from 1 to 4, most of them were observed with clear morning and evening peaks, which can be produced from O_3- and OH-initiated monoterpene oxidations. For the morning peak at around 7 am, the relative roles of O_3- and OH-initiated monoterpene oxidation were evaluated using the similar method as in Eq. (1) and Eq. (2). The reaction rate coefficient of monoterpene + OH is approximately 10^6 times higher than that of monoterpene + O_3 (Atkinson et al., 1990; Khamaganov and Hites, 2001; Gill and Hites, 2002; Hakola et al., 2012). In the morning, typical tropospheric OH concentrations have been observed to be around $1 \times 10^5 - 1 \times 10^6$ molecule cm^{-3} (0.004 − 0.04 ppt) (Shirinzadeh et al., 1987; Ren et al., 2003; Khan et al., 2008; Petäjä et al., 2009; Stone et al., 2012). For an OH concentration of 1×10^5 molecule cm^{-3} (0.004 ppt), with the average O_3 concentration of 15 ppb at 7 am, the reaction rate of monoterpene + OH was about 0.25 times as high as that of monoterpene + O_3. If the OH concentration was up to 1×10^6 molecule cm^{-3} (0.04 ppt) at 7 am, the reaction rate of monoterpene with OH radical was 2.5 times higher than that of monoterpene with O_3 according to the calculations. In other words, both oxidants are likely to be of importance at this time. For the evening peak of the less oxidized monoterpene oxidation products at 8 pm, the relative importance of O_3 and OH radical in monoterpene chemistry changed due to the lower OH concentration. With the average O_3 concentration of ~20 ppb and OH concentration of 0.012 ppt at 8 pm, the reaction rates of monoterpenes with O_3 and OH radical were at the similar level. Compared to other compounds, the evening peak of $C_9H_{14}O$, $C_{10}H_{16}O$, $C_{10}H_{18}O$, and $C_{10}H_{18}O_2$ extended over midnight. $C_9H_{14}O$ has been found to be one of the main products formed in the ozonolysis reactions of monoterpenes (Atkinson and Arey, 2003). O_3-initiated oxidation with extremely high monoterpene levels might be responsible for the high concentration of $C_9H_{14}O$ at night. Camphor ($C_{10}H_{16}O$), linalool ($C_{10}H_{18}O$), and linalool oxide ($C_{10}H_{18}O_2$) can be emitted by leaves and flowers (Corchnoy et al., 1992; Lavy et al., 2002). Therefore, direct emissions from vegetation in the Landes forest may contribute to the high mixing ratios of these compounds during night. With strong photochemical oxidations during the day, the diurnal cycles of the more oxidized compounds were characterized with a broad daytime distribution peaking between 2:00 pm and 4:00 pm UTC.

To date the oxidation processes of sesquiterpenes have been rarely investigated despite its potential significance in new particle formation and SOA formation (Bonn and Moortgat, 2003; Winterhalter et al., 2009). In this study, various

sesquiterpene oxidation products were observed, mainly including $C_{14}H_{22}O_n$, $C_{15}H_{22}O_n$, and $C_{15}H_{24}O_n$ (n=1~6), providing the possibility to explore the oxidations of sesquiterpenes in the atmosphere. As shown in Fig. 9 and Fig. S13-14, with the increase of oxygen numbers, sesquiterpene oxidation products displayed similar variations in their diurnal profiles with monoterpene oxidation products. The less oxidized products with 1 to 3 oxygen peaked both in the morning and in the evening, and the more oxidized compounds had a broad presence throughout the day. These results indicate the similar oxidation processes of sesquiterpenes with monoterpenes in the Landes forest.

3.4.3 Terpene-derived organic nitrates

Organic nitrates have been shown to represent a large fraction of submicron aerosol nitrate at both urban and rural sites in Europe (Kiendler-Scharr et al., 2016). During daytime, the reaction of peroxy radicals with NO can lead to the formation of organic nitrates. At night, NO_3 radicals from the oxidation of NO_2 by O_3, can also react with unsaturated compounds mostly coming from BVOCs to generate organic nitrates (Ayres et al., 2015). In this study, the less oxidized organic nitrates from monoterpene oxidations presented a distinct morning peak at 7 am (Fig. 11; Fig. S17-18), which can come from O_3- and OH-initiated monoterpene oxidations in the presence of NO_x. In addition, both isoprene- and monoterpene-derived organic nitrates showed evening peaks at around 8 pm (Fig. 10, Fig. S15-16). Using monoterpenes as an example, the relative roles of O_3, OH radical, and NO_3 radical in the nighttime formation of monoterpene-derived organic nitrates were evaluated by calculating the corresponding reaction rate (R):

$$R_{MT+O3} = k_{MT+O3}[MT][O_3] \quad (3)$$

$$R_{MT+OH} = k_{MT+OH}[MT][OH] \quad (4)$$

$$R_{MT+NO3} = k_{MT+NO3}[MT][NO_3] \quad (5)$$

where *k* is the reaction rate coefficient of monoterpenes with O_3, OH radical or NO_3 radical, and [MT], [O_3], [OH] or [NO_3] is the concentration of monoterpenes, O_3, OH radical or NO_3 radical.

Taking the peak concentration of monoterpene-derived organic nitrates at 8 pm as an example, the concentration of NO_3 radical was calculated by assuming a steady state between its production from O_3 and NO_2 and its removal by oxidation reactions and losses. The details have been described by Allan et al. (2000) and Peräkylä et al. (2014). With the high O_3 scavenging by monoterpenes in the evening, the estimated concentration of NO_3 radical was 0.017 ppt. Using k_{MT+O3} = 6.9×10^{-17} cm^3 $molecule^{-1}$ s^{-1} and k_{MT+NO3}= 7.5×10^{-12} cm^3 $molecule^{-1}$ s^{-1} taken from Peräkylä et al. (2014), the reaction rate of monoterpenes with O_3 was ~10 times higher than that of monoterpenes with NO_3 radicals. However, while ozonolysis was likely to dominate the overall oxidation of monoterpenes, the organic nitrate formation from O_3-initiated oxidation may still be much lower than those from NO_3-initiated oxidations, depending on what fraction of RO_2 radicals were reacting with NO_x. The relative importance of O_3 and OH radical in monoterpene chemistry at this time was the same as discussed in Sect. 3.4.2.

4. Conclusions

This work presented the deployment of the new state-of-the-art Vocus PTR-TOF in the French Landes forest during the CERVOLAND campaign. The Vocus PTR-TOF capabilities are evaluated for the first time in the actual ambient environment by the identification of the observed gas-phase molecules. With the improved detection efficiency and measurement precision compared to the conventional PTR instruments, multiple hydrocarbons with carbon numbers varying from 3 to 20 were observed as well as various VOCs oxidation products. Hydrocarbon signals were dominated by monoterpenes and their major fragment ions (e.g., $C_6H_8H^+$) within the instrument, consistent with high monoterpene emissions in the Landes forest. In general, most hydrocarbon molecules and the less oxidized compounds were characterized with high signals at night, whereas the more oxidized compounds exhibited elevated intensity during the day.

To demonstrate the importance of Vocus PTR-TOF application in atmospheric science study, the characteristics of terpenes and their oxidation products were investigated. In addition to the observation of isoprene, monoterpenes, and sesquiterpenes, this study presented the ambient characteristics of the rarely recorded diterpenes, which are traditionally considered as non-volatile species in the atmosphere. On average, the concentration of diterpenes was 1.7 ppt in the Landes forest, which was hundred to thousand times lower than that of monoterpenes (6.0 ppb) and sesquiterpenes (64.5 ppt). However, considering their low vapor pressure and high reactivity, diterpenes may potentially play an important part in atmospheric chemistry. The diurnal variations of diterpenes showed the maximum peak at night and low levels during the day, similar to those of monoterpenes and sesquiterpenes.

With strong photochemical oxidations of terpenes during the day, the more oxidized terpene reaction products were observed with a broad daytime peak, whereas the less oxidized terpene reaction products showed peak concentrations in the early morning or/and in the evening. By calculating the reaction rates of terpenes with the main oxidants, OH radical, O_3, and NO_3 radical, the contributions of different formation pathways to terpene oxidations were evaluated. The morning and evening peaks of non-nitrate terpene reaction products were contributed by both O_3- and OH-induced terpene oxidations. For the formation of terpene-derived organic nitrates, the relative importance of O_3-, OH-, and NO_3- driven oxidation pathways were more difficult to evaluate. Overall, we have shown that the Vocus PTR-TOF is able to detect a very broad coverage of compounds, from VOCs precursors to various oxidation products. Therefore, the application of the Vocus PTR-TOF in atmospheric sciences will be fundamental in understanding the chemical evolution of VOCs in the atmosphere and their roles in air quality and climate issues.

Author contributions

ME and MR conceived the study. MR, LH, PF, EV, and EP conducted the field measurements. HL carried out the data analysis. MR, PR, KD, JK, DW, MK, ME, and FB participated the data analysis. HL wrote the paper with inputs from all coauthors.

Competing interests

The authors declare that they have no conflict of interest.

Acknowledgements

This work was supported by the European Research Council under grants 742206 ATM-GP, 638703 COALA and 850614 CHAPAs, and the Academy of Finland (project numbers 317380 and 320094). The authors would like to thank the PRIMEQUAL programme for financial support (ADEME, convention #1662C0024). This study has also been carried out with financial support from the French National Research Agency (ANR) in the frame of the Investments for the future Programme, within the Cluster of Excellence COTE (ANR-10-LABX-45) of the University of Bordeaux. Special thanks to Dr Elena Ormeño-Lafuente (IMBE) for the loan of the BVOC calibration gas cylinders and Dr Christophe Chipeaux and Dr Denis Loustau (ISPA-INRA) for their precious help in providing meteorological data and access to ICOS station facility.

References

Atkinson, R., Hasegawa, D., and Aschmann, S. M.: Rate constants for the gas-phase reactions of O3 with a series of monoterpenes and related compounds at 296 ± 2 K, International Journal of Chemical Kinetics, 22, 871-887, 10.1002/kin.550220807, 1990.

Atkinson, R., and Arey, J.: Atmospheric Degradation of Volatile Organic Compounds, Chem Rev, 103, 4605-4638, 10.1021/cr0206420, 2003.

Ayres, B. R., Allen, H. M., Draper, D. C., Brown, S. S., Wild, R. J., Jimenez, J. L., Day, D. A., Campuzano-Jost, P., Hu, W., de Gouw, J., Koss, A., Cohen, R. C., Duffey, K. C., Romer, P., Baumann, K., Edgerton, E., Takahama, S., Thornton, J. A., Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Wennberg, P. O., Nguyen, T. B., Teng, A., Goldstein, A. H., Olson, K., and Fry, J. L.: Organic nitrate aerosol formation via NO₃ + biogenic volatile organic compounds in the southeastern United States, Atmos. Chem. Phys., 15, 13377-13392, 10.5194/acp-15-13377-2015, 2015.

Baltensperger, U., Kalberer, M., Dommen, J., Paulsen, D., Alfarra, M. R., Coe, H., Fisseha, R., Gascho, A., Gysel, M., Nyeki, S., Sax, M., Steinbacher, M., Prevot, A. S. H., Sjögren, S., Weingartner, E., and Zenobi, R.: Secondary organic aerosols from anthropogenic and biogenic precursors, Faraday Discussions, 130, 265-278, 10.1039/B417367H, 2005.

Bernhammer, A. K., Fischer, L., Mentler, B., Heinritzi, M., Simon, M., and Hansel, A.: Production of highly oxygenated organic molecules (HOMs) from trace contaminants during isoprene oxidation, Atmos. Meas. Tech., 11, 4763-4773, 10.5194/amt-11-4763-2018, 2018.

Bertram, T. H., Kimmel, J. R., Crisp, T. A., Ryder, O. S., Yatavelli, R. L. N., Thornton, J. A., Cubison, M. J., Gonin, M., and Worsnop, D. R.: A field-deployable, chemical ionization time-of-flight mass spectrometer, Atmos. Meas. Tech., 4, 1471-1479, 10.5194/amt-4-1471-2011, 2011.

Bianchi, F., Kurtén, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crounse, J. D., Wennberg, P. O., Mentel, T. F., Wildt, J., Junninen, H., Jokinen, T., Kulmala, M., Worsnop, D. R., Thornton, J. A., Donahue, N., Kjaergaard, H. G., and Ehn, M.: Highly Oxygenated Organic Molecules (HOM) from Gas-Phase Autoxidation Involving Peroxy Radicals: A Key Contributor to Atmospheric Aerosol, Chem Rev, 10.1021/acs.chemrev.8b00395, 2019.

Bonn, B., and Moortgat, G. K.: Sesquiterpene ozonolysis: Origin of atmospheric new particle formation from biogenic hydrocarbons, 30, doi:10.1029/2003GL017000, 2003.

Boy, M., Bonn, B., Maso, M. D., Hakola, H., Hirsikko, A., Kulmala, M., Kurtén, T., Laakso, L., Mäkelä, J., Riipinen, I., Rannik, Ü., Sihto, S.-L., and Ruuskanen, T. M.: Biogenic Sesquiterpenes and Atmospheric New Particle Formation: A Boreal Forest Site Investigation, Nucleation and Atmospheric Aerosols, Dordrecht, 2007, 344-349,

Boyd, C. M., Sanchez, J., Xu, L., Eugene, A. J., Nah, T., Tuet, W. Y., Guzman, M. I., and Ng, N. L.: Secondary organic aerosol formation from the β-pinene + NO3 system: effect of humidity and peroxy radical fate, Atmos. Chem. Phys., 15, 7497-7522, 10.5194/acp-15-7497-2015, 2015.

Breitenlechner, M., Fischer, L., Hainer, M., Heinritzi, M., Curtius, J., and Hansel, A.: PTR3: An Instrument for Studying the Lifecycle of Reactive Organic Carbon in the Atmosphere, Analytical Chemistry, 89, 5824-5831, 10.1021/acs.analchem.6b05110, 2017.

Bsaibes, S., Al Ajami, M., Mermet, K., Truong, F., Batut, S., Hecquet, C., Dusanter, S., Léornadis, T., Sauvage, S., Kammer, J., Flaud, P.-M., Perraudin, E., Villenave, E., Locoge, N., Gros, V., and Schoemaecker, C.: Variability of OH reactivity in the Landes maritime Pine forest: Results from the LANDEX campaign 2017, Atmos. Chem. Phys. Discuss., https://doi.org/10.5194/acp-2019-548, in review, 2019.

Buhr, K., van Ruth, S., and Delahunty, C.: Analysis of volatile flavour compounds by Proton Transfer Reaction-Mass Spectrometry: fragmentation patterns and discrimination between isobaric and isomeric compounds, International Journal of Mass Spectrometry, 221, 1-7, https://doi.org/10.1016/S1387-3806(02)00896-5, 2002.

Buser, A. M., Kierkegaard, A., Bogdal, C., MacLeod, M., Scheringer, M., and Hungerbühler, K.: Concentrations in Ambient Air and Emissions of Cyclic Volatile Methylsiloxanes in Zurich, Switzerland, Environmental Science & Technology, 47, 7045-7051, 10.1021/es3046586, 2013.

Calfapietra, C., Fares, S., Manes, F., Morani, A., Sgrigna, G., and Loreto, F.: Role of Biogenic Volatile Organic Compounds (BVOC) emitted by urban trees on ozone concentration in cities: A review, Environ Pollut, 183, 71-80, https://doi.org/10.1016/j.envpol.2013.03.012, 2013.

Carlton, A. G., Wiedinmyer, C., and Kroll, J. H.: A review of Secondary Organic Aerosol (SOA) formation from isoprene, Atmos. Chem. Phys., 9, 4987-5005, 10.5194/acp-9-4987-2009, 2009.

Carslaw, K. S., Lee, L. A., Reddington, C. L., Pringle, K. J., Rap, A., Forster, P. M., Mann, G. W., Spracklen, D. V., Woodhouse, M. T., Regayre, L. A., and Pierce, J. R.: Large contribution of natural aerosols to uncertainty in indirect forcing, Nature, 503, 67, 10.1038/nature12674, 2013.

Claeys, M., Graham, B., Vas, G., Wang, W., Vermeylen, R., Pashynska, V., Cafmeyer, J., Guyon, P., Andreae, M. O., Artaxo, P., and Maenhaut, W.: Formation of Secondary Organic Aerosols Through Photooxidation of Isoprene, 303, 1173-1176, 10.1126/science.1092805 %J Science, 2004.

Corchnoy, S. B., Arey, J., and Atkinson, R.: Hydrocarbon emissions from twelve urban shade trees of the Los Angeles, California, Air Basin, Atmospheric Environment. Part B. Urban Atmosphere, 26, 339-348, https://doi.org/10.1016/0957-1272(92)90009-H, 1992.

Criegee, R.: Mechanism of Ozonolysis, Angewandte Chemie International Edition in English, 14, 745-752, 10.1002/anie.197507451, 1975.

Crounse, J. D., McKinney, K. A., Kwan, A. J., and Wennberg, P. O.: Measurement of Gas-Phase Hydroperoxides by Chemical Ionization Mass Spectrometry, Analytical Chemistry, 78, 6726-6732, 10.1021/ac0604235, 2006.

Donahue, N. M., Kroll, J. H., Pandis, S. N., and Robinson, A. L.: A two-dimensional volatility basis set – Part 2: Diagnostics of organic-aerosol evolution, Atmos. Chem. Phys., 12, 615-634, 10.5194/acp-12-615-2012, 2012.

Dreyfus, G. B., Schade, G. W., and Goldstein, A. H.: Observational constraints on the contribution of isoprene oxidation to ozone production on the western slope of the Sierra Nevada, California, Journal of Geophysical Research: Atmospheres, 107, ACH 1-1-ACH 1-17, 10.1029/2001JD001490, 2002.

Dusanter, S., Vimal, D., and Stevens, P. S.: Technical note: Measuring tropospheric OH and HO2 by laser-induced fluorescence at low pressure. A comparison of calibration techniques, Atmos. Chem. Phys., 8, 321–340, https://doi.org/10.5194/acp-8-321-2008, 2008.

Ehn, M., Kleist, E., Junninen, H., Petäjä, T., Lönn, G., Schobesberger, S., Dal Maso, M., Trimborn, A., Kulmala, M., Worsnop, D. R., Wahner, A., Wildt, J., and Mentel, T. F.: Gas phase formation of extremely oxidized pinene reaction products in chamber and ambient air, Atmos. Chem. Phys., 12, 5113-5127, 10.5194/acp-12-5113-2012, 2012.

Ehn, M., Thornton, J. A., Kleist, E., Sipila, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I. H., Rissanen, M., Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurten, T., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., Canagaratna, M., Dal Maso, M., Berndt, T., Petaja, T., Wahner, A., Kerminen, V. M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol, Nature, 506, 476-+, 2014.

Gill, K. J., and Hites, R. A.: Rate Constants for the Gas-Phase Reactions of the Hydroxyl Radical with Isoprene, α- and β-Pinene, and Limonene as a Function of Temperature, The Journal of Physical Chemistry A, 106, 2538-2544, 10.1021/jp013532q, 2002.

Goldstein, A. H., McKay, M., Kurpius, M. R., Schade, G. W., Lee, A., Holzinger, R., and Rasmussen, R. A.: Forest thinning experiment confirms ozone deposition to forest canopy is dominated by reaction with biogenic VOCs, Geophys Res Lett, 31, 10.1029/2004GL021259, 2004.

Gueneron, M., Erickson, M. H., VanderSchelden, G. S., and Jobson, B. T.: PTR-MS fragmentation patterns of gasoline hydrocarbons, International Journal of Mass Spectrometry, 379, 97-109, https://doi.org/10.1016/j.ijms.2015.01.001, 2015.

Guenther, A., Hewitt, C. N., Erickson, D., Fall, R., Geron, C., Graedel, T., Harley, P., Klinger, L., Lerdau, M., McKay, W. A., Pierce, T., Scholes, B., Steinbrecher, R., Tallamraju, R., Taylor, J., and Zimmerman, P.: A global model of natural volatile organic compound emissions, Journal of Geophysical Research: Atmospheres, 100, 8873-8892, 10.1029/94JD02950, 1995.

Hakola, H., Tarvainen, V., Bäck, J., Ranta, H., Bonn, B., Rinne, J., and Kulmala, M.: Seasonal variation of mono- and sesquiterpene emission rates of Scots pine, Biogeosciences, 3, 93-101, 10.5194/bg-3-93-2006, 2006.

Hakola, H., Hellén, H., Hemmilä, M., Rinne, J., and Kulmala, M.: In situ measurements of volatile organic compounds in a boreal forest, Atmos. Chem. Phys., 12, 11665-11678, 10.5194/acp-12-11665-2012, 2012.

Hallquist, M., Wängberg, I., Ljungström, E., Barnes, I., and Becker, K.-H.: Aerosol and Product Yields from NO3 Radical-Initiated Oxidation of Selected Monoterpenes, Environmental Science & Technology, 33, 553-559, 10.1021/es980292s, 1999.

Hatfield, M. L., and Huff Hartz, K. E.: Secondary organic aerosol from biogenic volatile organic compound mixtures, Atmos Environ, 45, 2211-2219, https://doi.org/10.1016/j.atmosenv.2011.01.065, 2011.

Hellén, H., Hakola, H., Pystynen, K.-H., Rinne, J., and Haapanala, S.: C2-C10 hydrocarbon emissions from a boreal wetland and forest floor, Biogeosciences, 3, 167–174, https://doi.org/10.5194/bg-3-167-2006, 2006.

Hellén, H., Praplan, A. P., Tykkä, T., Ylivinkka, I., Vakkari, V., Bäck, J., Petäjä, T., Kulmala, M., and Hakola, H.: Long-term measurements of volatile organic compounds highlight the importance of sesquiterpenes for the atmospheric chemistry of a boreal forest, Atmos. Chem. Phys., 18, 13839-13863, 10.5194/acp-18-13839-2018, 2018.

548 Henze, D. K., and Seinfeld, J. H.: Global secondary organic aerosol from isoprene oxidation, Geophys Res Lett, 33,
549 10.1029/2006GL025976, 2006.

550 Holzinger, R., Acton, W. J. F., Bloss, W. J., Breitenlechner, M., Crilley, L. R., Dusanter, S., Gonin, M., Gros, V., Keutsch, F.
551 N., Kiendler-Scharr, A., Kramer, L. J., Krechmer, J. E., Languille, B., Locoge, N., Lopez-Hilfiker, F., Materić, D., Moreno,
552 S., Nemitz, E., Quéléver, L. L. J., Sarda Esteve, R., Sauvage, S., Schallhart, S., Sommariva, R., Tillmann, R., Wedel, S.,
553 Worton, D. R., Xu, K., and Zaytsev, A.: Validity and limitations of simple reaction kinetics to calculate concentrations of
554 organic compounds from ion counts in PTR-MS, Atmos. Meas. Tech., 12, 6193–6208, https://doi.org/10.5194/amt-12-6193-
555 2019, 2019.

556 Hyttinen, N., Kupiainen-Määttä, O., Rissanen, M. P., Muuronen, M., Ehn, M., and Kurtén, T.: Modeling the Charging of
557 Highly Oxidized Cyclohexene Ozonolysis Products Using Nitrate-Based Chemical Ionization, The Journal of Physical
558 Chemistry A, 119, 6339-6345, 10.1021/acs.jpca.5b01818, 2015.

559 Intergovernmental Panel on Climate Change (IPCC), Climate Change 2013: The Physical Science Basis. Contribution of
560 Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, 1535 pp., Cambridge
561 Univ. Press, Cambridge, U. K., and New York, 2013.

562 Isaacman-VanWertz, G., Massoli, P., O'Brien, R., Lim, C., Franklin, J. P., Moss, J. A., Hunter, J. F., Nowak, J. B., Canagaratna,
563 M. R., Misztal, P. K., Arata, C., Roscioli, J. R., Herndon, S. T., Onasch, T. B., Lambe, A. T., Jayne, J. T., Su, L. P., Knopf, D.
564 A., Goldstein, A. H., Worsnop, D. R., and Kroll, J. H.: Chemical evolution of atmospheric organic carbon over multiple
565 generations of oxidation, Nat Chem, 10, 462-468, 2018.

566 Jardine, K., Yañez Serrano, A., Arneth, A., Abrell, L., Jardine, A., van Haren, J., Artaxo, P., Rizzo, L. V., Ishida, F. Y., Karl,
567 T., Kesselmeier, J., Saleska, S., and Huxman, T.: Within-canopy sesquiterpene ozonolysis in Amazonia, 116,
568 doi:10.1029/2011JD016243, 2011.

569 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe,
570 H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson,
571 K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M.,
572 Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R.,
573 Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D.,
574 Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel,
575 J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C.
576 E., Baltensperger, U., and Worsnop, D. R.: Evolution of Organic Aerosols in the Atmosphere, 326, 1525-1529,
577 10.1126/science.1180353 %J Science, 2009.

578 Jokinen, T., Sipilä, M., Richters, S., Kerminen, V.-M., Paasonen, P., Stratmann, F., Worsnop, D., Kulmala, M., Ehn, M.,
579 Herrmann, H., and Berndt, T.: Rapid Autoxidation Forms Highly Oxidized RO2 Radicals in the Atmosphere, 53, 14596-14600,
580 doi:10.1002/anie.201408566, 2014.

581 Jokinen, T., Berndt, T., Makkonen, R., Kerminen, V.-M., Junninen, H., Paasonen, P., Stratmann, F., Herrmann, H., Guenther,
582 A. B., Worsnop, D. R., Kulmala, M., Ehn, M., and Sipilä, M.: Production of extremely low volatile organic compounds from
583 biogenic emissions: Measured yields and atmospheric implications, 112, 7123-7128, 10.1073/pnas.1423977112 %J
584 Proceedings of the National Academy of Sciences, 2015.

585 Kammer, J., Perraudin, E., Flaud, P. M., Lamaud, E., Bonnefond, J. M., and Villenave, E.: Observation of nighttime new
586 particle formation over the French Landes forest, Sci Total Environ, 621, 1084-1092,
587 https://doi.org/10.1016/j.scitotenv.2017.10.118, 2018.

588 Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., Van Dingenen, R., Ervens, B., Nenes,
589 A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski, Y., Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter, R., Myhre,
590 C. E. L., Tsigaridis, K., Vignati, E., Stephanou, E. G., and Wilson, J.: Organic aerosol and global climate modelling: a review,
591 Atmos. Chem. Phys., 5, 1053–1123, https://doi.org/10.5194/acp-5-1053-2005, 2005.

592 Kanawade, V. P., Jobson, B. T., Guenther, A. B., Erupe, M. E., Pressley, S. N., Tripathi, S. N., and Lee, S. H.: Isoprene
593 suppression of new particle formation in a mixed deciduous forest, Atmos. Chem. Phys., 11, 6013-6027, 10.5194/acp-11-
594 6013-2011, 2011.

595 Karl, M., Brauers, T., Dorn, H. P., Holland, F., Komenda, M., Poppe, D., Rohrer, F., Rupp, L., Schaub, A., and Wahner, A.:
596 Kinetic Study of the OH-isoprene and O3-isoprene reaction in the atmosphere simulation chamber, SAPHIR, Geophys Res
597 Lett, 31, 10.1029/2003GL019189, 2004.

Karl, T., Hansel, A., Cappellin, L., Kaser, L., Herdlinger-Blatt, I., and Jud, W.: Selective measurements of isoprene and 2-methyl-3-buten-2-ol based on NO+ ionization mass spectrometry, Atmos. Chem. Phys., 12, 11877–11884, https://doi.org/10.5194/acp-12-11877-2012, 2012.

Kari, E., Miettinen, P., Yli-Pirilä, P., Virtanen, A., and Faiola, C. L.: PTR-ToF-MS product ion distributions and humidity-dependence of biogenic volatile organic compounds, International Journal of Mass Spectrometry, 430, 87-97, https://doi.org/10.1016/j.ijms.2018.05.003, 2018.

Kaser, L., Karl, T., Guenther, A., Graus, M., Schnitzhofer, R., Turnipseed, A., Fischer, L., Harley, P., Madronich, M., Gochis, D., Keutsch, F. N., and Hansel, A.: Undisturbed and disturbed above canopy ponderosa pine emissions: PTR-TOF-MS measurements and MEGAN 2.1 model results, Atmos. Chem. Phys., 13, 11935–11947, https://doi.org/10.5194/acp-13-11935-2013, 2013.

Keck, L., Hoeschen, C., and Oeh, U.: Effects of carbon dioxide in breath gas on proton transfer reaction-mass spectrometry (PTR-MS) measurements, International Journal of Mass Spectrometry, 270, 156-165, https://doi.org/10.1016/j.ijms.2007.12.009, 2008.

Keeling, C. I., and Bohlmann, J.: Diterpene resin acids in conifers, Phytochemistry, 67, 2415-2423, https://doi.org/10.1016/j.phytochem.2006.08.019, 2006.

Kesselmeier, J., and Staudt, M.: Biogenic Volatile Organic Compounds (VOC): An Overview on Emission, Physiology and Ecology, J Atmos Chem, 33, 23-88, 10.1023/A:1006127516791, 1999.

Khamaganov, V. G., and Hites, R. A.: Rate Constants for the Gas-Phase Reactions of Ozone with Isoprene, α- and β-Pinene, and Limonene as a Function of Temperature, The Journal of Physical Chemistry A, 105, 815-822, 10.1021/jp002730z, 2001.

Khan, M. A. H., Ashfold, M. J., Nickless, G., Martin, D., Watson, L. A., Hamer, P. D., Wayne, R. P., Canosa-Mas, C. E., and Shallcross, D. E.: Night-time NO3 and OH radical concentrations in the United Kingdom inferred from hydrocarbon measurements, Atmospheric Science Letters, 9, 140-146, 10.1002/asl.175, 2008.

Kiendler-Scharr, A., Wildt, J., Maso, M. D., Hohaus, T., Kleist, E., Mentel, T. F., Tillmann, R., Uerlings, R., Schurr, U., and Wahner, A.: New particle formation in forests inhibited by isoprene emissions, Nature, 461, 381, 10.1038/nature08292 https://www.nature.com/articles/nature08292#supplementary-information, 2009.

Kiendler-Scharr, A., Mensah, A. A., Friese, E., Topping, D., Nemitz, E., Prevot, A. S. H., Äijälä, M., Allan, J., Canonaco, F., Canagaratna, M., Carbone, S., Crippa, M., Dall Osto, M., Day, D. A., De Carlo, P., Di Marco, C. F., Elbern, H., Eriksson, A., Freney, E., Hao, L., Herrmann, H., Hildebrandt, L., Hillamo, R., Jimenez, J. L., Laaksonen, A., McFiggans, G., Mohr, C., O'Dowd, C., Otjes, R., Ovadnevaite, J., Pandis, S. N., Poulain, L., Schlag, P., Sellegri, K., Swietlicki, E., Tiitta, P., Vermeulen, A., Wahner, A., Worsnop, D., and Wu, H. C.: Ubiquity of organic nitrates from nighttime chemistry in the European submicron aerosol, Geophys Res Lett, 43, 7735-7744, 10.1002/2016GL069239, 2016.

Kim, S., Karl, T., Helmig, D., Daly, R., Rasmussen, R., and Guenther, A.: Measurement of atmospheric sesquiterpenes by proton transfer reaction-mass spectrometry (PTR-MS), Atmos. Meas. Tech., 2, 99-112, 10.5194/amt-2-99-2009, 2009.

Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M., Yan, C., Almeida, J., Tröstl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim, A., Bernhammer, A.-K., Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X., Craven, J., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Kim, J., Krapf, M., Kürten, A., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Molteni, U., Onnela, A., Peräkylä, O., Piel, F., Petäjä, T., Praplan, A. P., Pringle, K., Rap, A., Richards, N. A. D., Riipinen, I., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Scott, C. E., Seinfeld, J. H., Sipilä, M., Steiner, G., Stozhkov, Y., Stratmann, F., Tomé, A., Virtanen, A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R., Baltensperger, U., Kulmala, M., Carslaw, K. S., and Curtius, J.: Ion-induced nucleation of pure biogenic particles, Nature, 533, 521, 10.1038/nature17953, 2016.

Krechmer, J., Lopez-Hilfiker, F., Koss, A., Hutterli, M., Stoermer, C., Deming, B., Kimmel, J., Warneke, C., Holzinger, R., Jayne, J., Worsnop, D., Fuhrer, K., Gonin, M., and de Gouw, J.: Evaluation of a New Reagent-Ion Source and Focusing Ion–Molecule Reactor for Use in Proton-Transfer-Reaction Mass Spectrometry, Analytical Chemistry, 90, 12011-12018, 10.1021/acs.analchem.8b02641, 2018.

Lavy, M., Zuker, A., Lewinsohn, E., Larkov, O., Ravid, U., Vainstein, A., and Weiss, D.: Linalool and linalool oxide production in transgenic carnation flowers expressing the Clarkia breweri linalool synthase gene, Molecular Breeding, 9, 103-111, 10.1023/A:1026755414773, 2002.

Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Kurten, T., Worsnop, D. R., and Thornton, J. A.: An Iodide-Adduct High-Resolution Time-of-Flight Chemical-Ionization Mass Spectrometer: Application to Atmospheric Inorganic and Organic Compounds, Environmental Science & Technology, 48, 6309-6317, 2014.

Lide, D. R. CRC Handbook of Chemistry and Physics; Internet Version 2005, <http://www.hbcpnetbase.com>, CRC Press, Boca Raton, FL, 2005.

Liu, X., Deming, B., Pagonis, D., Day, D. A., Palm, B. B., Talukdar, R., Roberts, J. M., Veres, P. R., Krechmer, J. E., Thornton, J. A., de Gouw, J. A., Ziemann, P. J., and Jimenez, J. L.: Effects of Gas-Wall Interactions on Measurements of Semivolatile Compounds and Small Polar Molecules, Atmos. Meas. Tech. Discuss., https://doi.org/10.5194/amt-2019-52, in review, 2019.

Maleknia, S. D., Bell, T. L., and Adams, M. A.: PTR-MS analysis of reference and plant-emitted volatile organic compounds, International Journal of Mass Spectrometry, 262, 203-210, https://doi.org/10.1016/j.ijms.2006.11.010, 2007.

Maria, S. F., Russell, L. M., Gilles, M. K., and Myneni, S. C. B.: Organic Aerosol Growth Mechanisms and Their Climate-Forcing Implications, 306, 1921-1924, 10.1126/science.1103491 %J Science, 2004.

Massoli, P., Stark, H., Canagaratna, M. R., Krechmer, J. E., Xu, L., Ng, N. L., Mauldin, R. L., Yan, C., Kimmel, J., Misztal, P. K., Jimenez, J. L., Jayne, J. T., and Worsnop, D. R.: Ambient Measurements of Highly Oxidized Gas-Phase Molecules during the Southern Oxidant and Aerosol Study (SOAS) 2013, ACS Earth and Space Chemistry, 2, 653-672, 10.1021/acsearthspacechem.8b00028, 2018.

Matsunaga, S. N., Chatani, S., Nakatsuka, S., Kusumoto, D., Kubota, K., Utsumi, Y., Enoki, T., Tani, A., and Hiura, T.: Determination and potential importance of diterpene (kaur-16-ene) emitted from dominant coniferous trees in Japan, Chemosphere, 87, 886-893, https://doi.org/10.1016/j.chemosphere.2012.01.040, 2012.

Mauderly, J. L., and Chow, J. C.: Health Effects of Organic Aerosols, Inhalation Toxicology, 20, 257-288, 10.1080/08958370701866008, 2008.

McFiggans, G., Mentel, T. F., Wildt, J., Pullinen, I., Kang, S., Kleist, E., Schmitt, S., Springer, M., Tillmann, R., Wu, C., Zhao, D., Hallquist, M., Faxon, C., Le Breton, M., Hallquist, Å. M., Simpson, D., Bergström, R., Jenkin, M. E., Ehn, M., Thornton, J. A., Alfarra, M. R., Bannan, T. J., Percival, C. J., Priestley, M., Topping, D., and Kiendler-Scharr, A.: Secondary organic aerosol reduced by mixture of atmospheric vapours, Nature, 565, 587-593, 10.1038/s41586-018-0871-y, 2019.

Mermet, K., Sauvage, S., Dusanter, S., Salameh, T., Léonardis, T., Flaud, P.-M., Perraudin, É., Villenave, É., and Locoge, N.: Optimization of a gas chromatographic unit for measuring BVOCs in ambient air, Atmos. Meas. Tech. Discuss., https://doi.org/10.5194/amt-2019-224, in review, 2019.

Monson, R. K., and Fall, R.: Isoprene Emission from Aspen Leaves, Influence of Environment and Relation to Photosynthesis and Photorespiration, 90, 267-274, 10.1104/pp.90.1.267 %J Plant Physiology, 1989.

Moreaux, V., Lamaud, É., Bosc, A., Bonnefond, J.-M., Medlyn, B. E., and Loustau, D.: Paired comparison of water, energy and carbon exchanges over two young maritime pine stands (Pinus pinaster Ait.): effects of thinning and weeding in the early stage of tree growth, Tree Physiology, 31, 903-921, 10.1093/treephys/tpr048, 2011.

Ng, N. L., Brown, S. S., Archibald, A. T., Atlas, E., Cohen, R. C., Crowley, J. N., Day, D. A., Donahue, N. M., Fry, J. L., Fuchs, H., Griffin, R. J., Guzman, M. I., Herrmann, H., Hodzic, A., Iinuma, Y., Jimenez, J. L., Kiendler-Scharr, A., Lee, B. H., Luecken, D. J., Mao, J., McLaren, R., Mutzel, A., Osthoff, H. D., Ouyang, B., Picquet-Varrault, B., Platt, U., Pye, H. O. T., Rudich, Y., Schwantes, R. H., Shiraiwa, M., Stutz, J., Thornton, J. A., Tilgner, A., Williams, B. J., and Zaveri, R. A.: Nitrate radicals and biogenic volatile organic compounds: oxidation, mechanisms, and organic aerosol, Atmos. Chem. Phys., 17, 2103-2162, 10.5194/acp-17-2103-2017, 2017.

Peräkylä, O., Vogt, M., Tikkanen, O.-P., Laurila, T., Kajos, M. K., Rantala, P. A., Patokoski, J., Aalto, J., Yli-Juuti, T., Ehn, M., Sipilä, M., Paasonen, P., Rissanen, M., Nieminen, T., Taipale, R., Keronen, P., Lappalainen, H. K., Ruuskanen, T. M., Rinne, J., Kerminen, V.-M., Kulmala, M., Bäck, J., and Petäjä, T.: Monoterpenes' oxidation capacity and rate over a boreal forest: temporal variation and connection to growth of newly formed particles, Boreal Environ. Res., 19, 293–310, 2014.

Petäjä, T., Mauldin, I. R. L., Kosciuch, E., McGrath, J., Nieminen, T., Paasonen, P., Boy, M., Adamov, A., Kotiaho, T., and Kulmala, M.: Sulfuric acid and OH concentrations in a boreal forest site, Atmos. Chem. Phys., 9, 7435-7448, 10.5194/acp-9-7435-2009, 2009.

Pfeiffer, T., Forberich, O., and Comes, F. J.: The contribution of the ozonolysis of terpenes to tropospheric OH concentrations, Canadian Journal of Physics, 79, 131-142, 10.1139/p01-030, 2001.

Presto, A. A., Huff Hartz, K. E., and Donahue, N. M.: Secondary Organic Aerosol Production from Terpene Ozonolysis. 2. Effect of NOx Concentration, Environmental Science & Technology, 39, 7046-7054, 10.1021/es050400s, 2005.

Qin, M., Hu, Y., Wang, X., Vasilakos, P., Boyd, C. M., Xu, L., Song, Y., Ng, N. L., Nenes, A., and Russell, A. G.: Modeling biogenic secondary organic aerosol (BSOA) formation from monoterpene reactions with NO3: A case study of the SOAS campaign using CMAQ, Atmos Environ, 184, 146-155, https://doi.org/10.1016/j.atmosenv.2018.03.042, 2018.

Ren, X., Harder, H., Martinez, M., Lesher, R. L., Oliger, A., Shirley, T., Adams, J., Simpas, J. B., and Brune, W. H.: HOx concentrations and OH reactivity observations in New York City during PMTACS-NY2001, Atmos Environ, 37, 3627-3637, https://doi.org/10.1016/S1352-2310(03)00460-6, 2003.

Rickard, A. R., Johnson, D., McGill, C. D., and Marston, G.: OH yields in the gas-phase reactions of ozone with alkenes, J Phys Chem A, 103, 7656-7664, 1999.

Riva, M., Rantala, P., Krechmer, J. E., Peräkylä, O., Zhang, Y., Heikkinen, L., Garmash, O., Yan, C., Kulmala, M., Worsnop, D., and Ehn, M.: Evaluating the performance of five different chemical ionization techniques for detecting gaseous oxygenated organic species, Atmos. Meas. Tech., 12, 2403-2421, https://doi.org/10.5194/amt-12-2403-2019, 2019a.

Riva, M., Heikkinen, L., Bell, D. M., Peräkylä, O., Zha, Q., Schallhart, S., Rissanen, M. P., Imre, D., Petäjä, T., Thornton, J. A., Zelenyuk, A., and Ehn, M.: Chemical transformations in monoterpene-derived organic aerosol enhanced by inorganic composition, npj Climate and Atmospheric Science, 2, 2, 10.1038/s41612-018-0058-0, 2019b.

Ruppert, L., and Heinz Becker, K.: A product study of the OH radical-initiated oxidation of isoprene: formation of C5-unsaturated diols, Atmos Environ, 34, 1529-1542, https://doi.org/10.1016/S1352-2310(99)00408-2, 2000.

Sakulyanontvittaya, T., Duhl, T., Wiedinmyer, C., Helmig, D., Matsunaga, S., Potosnak, M., Milford, J., and Guenther, A.: Monoterpene and Sesquiterpene Emission Estimates for the United States, Environmental Science & Technology, 42, 1623-1629, 10.1021/es702274e, 2008.

Schwarz, K., Filipiak, W., and Amann, A.: Determining concentration patterns of volatile compounds in exhaled breath by PTR-MS, Journal of Breath Research, 3, 027002, 10.1088/1752-7155/3/2/027002, 2009.

Sekimoto, K., Li, S.-M., Yuan, B., Koss, A., Coggon, M., Warneke, C., and de Gouw, J.: Calculation of the sensitivity of proton-transfer-reaction mass spectrometry (PTR-MS) for organic trace gases using molecular properties, International Journal of Mass Spectrometry, 421, 71-94, https://doi.org/10.1016/j.ijms.2017.04.006, 2017.

Shirinzadeh, B., Wang, C. C., and Deng, D. Q.: Diurnal variation of the OH concentration in ambient air, Geophys Res Lett, 14, 123-126, 10.1029/GL014i002p00123, 1987.

Simon, V., Clement, B., Riba, M. - L., and Torres, L.: The Landes experiment: Monoterpenes emitted from the maritime pine, J. Geophys. Res., 99(D8), 16501 – 16510, doi:10.1029/94JD00785, 1994.

Sindelarova, K., Granier, C., Bouarar, I., Guenther, A., Tilmes, S., Stavrakou, T., Müller, J. F., Kuhn, U., Stefani, P., and Knorr, W.: Global data set of biogenic VOC emissions calculated by the MEGAN model over the last 30 years, Atmos. Chem. Phys., 14, 9317-9341, 10.5194/acp-14-9317-2014, 2014.

Stone, D., Whalley, L. K., and Heard, D. E.: Tropospheric OH and HO2 radicals: field measurements and model comparisons, Chemical Society Reviews, 41, 6348-6404, 10.1039/C2CS35140D, 2012.

Surratt, J. D., Chan, A. W. H., Eddingsaas, N. C., Chan, M., Loza, C. L., Kwan, A. J., Hersey, S. P., Flagan, R. C., Wennberg, P. O., and Seinfeld, J. H.: Reactive intermediates revealed in secondary organic aerosol formation from isoprene, 107, 6640-6645, 10.1073/pnas.0911114107 %J Proceedings of the National Academy of Sciences, 2010.

Tani, A., Hayward, S., and Hewitt, C. N.: Measurement of monoterpenes and related compounds by proton transfer reaction-mass spectrometry (PTR-MS), International Journal of Mass Spectrometry, 223-224, 561-578, https://doi.org/10.1016/S1387-3806(02)00880-1, 2003.

Tani, A., Hayward, S., Hansel, A., and Hewitt, C. N.: Effect of water vapour pressure on monoterpene measurements using proton transfer reaction-mass spectrometry (PTR-MS), International Journal of Mass Spectrometry, 239, 161-169, https://doi.org/10.1016/j.ijms.2004.07.020, 2004.

Tani, A.: Fragmentation and Reaction Rate Constants of Terpenoids Determined by Proton Transfer Reaction-mass Spectrometry, Environmental Control in Biology, 51, 23-29, 10.2525/ecb.51.23, 2013.

Tsigaridis, K., and Kanakidou, M.: Global modelling of secondary organic aerosol in the troposphere: a sensitivity analysis, Atmos. Chem. Phys., 3, 1849-1869, 10.5194/acp-3-1849-2003, 2003.

von Schwartzenberg, K., Schultze, W., and Kassner, H. J. P. C. R.: The moss Physcomitrella patens releases a tetracyclic diterpene, 22, 780-786, 10.1007/s00299-004-0754-6, 2004.

Warneke, C., de Gouw, J. A., Kuster, W. C., Goldan, P. D., and Fall, R.: Validation of Atmospheric VOC Measurements by Proton-Transfer- Reaction Mass Spectrometry Using a Gas-Chromatographic Preseparation Method, Environmental Science & Technology, 37, 2494-2501, 10.1021/es026266i, 2003.

Wei, D., Fuentes, J. D., Gerken, T., Chamecki, M., Trowbridge, A. M., Stoy, P. C., Katul, G. G., Fisch, G., Acevedo, O., Manzi, A., von Randow, C., and dos Santos, R. M. N.: Environmental and biological controls on seasonal patterns of isoprene above a rain forest in central Amazonia, Agricultural and Forest Meteorology, 256-257, 391-406, https://doi.org/10.1016/j.agrformet.2018.03.024, 2018.

Wennberg, P. O., Bates, K. H., Crounse, J. D., Dodson, L. G., McVay, R. C., Mertens, L. A., Nguyen, T. B., Praske, E., Schwantes, R. H., Smarte, M. D., St Clair, J. M., Teng, A. P., Zhang, X., and Seinfeld, J. H.: Gas-Phase Reactions of Isoprene and Its Major Oxidation Products, Chem Rev, 118, 3337-3390, 2018.

Winterhalter, R., Herrmann, F., Kanawati, B., Nguyen, T. L., Peeters, J., Vereecken, L., and Moortgat, G. K.: The gas-phase ozonolysis of β-caryophyllene (C15H24). Part I: an experimental study, Physical Chemistry Chemical Physics, 11, 4152-4172, 10.1039/B817824K, 2009.

Xiong, F., McAvey, K. M., Pratt, K. A., Groff, C. J., Hostetler, M. A., Lipton, M. A., Starn, T. K., Seeley, J. V., Bertman, S. B., Teng, A. P., Crounse, J. D., Nguyen, T. B., Wennberg, P. O., Misztal, P. K., Goldstein, A. H., Guenther, A. B., Koss, A. R., Olson, K. F., de Gouw, J. A., Baumann, K., Edgerton, E. S., Feiner, P. A., Zhang, L., Miller, D. O., Brune, W. H., and Shepson, P. B.: Observation of isoprene hydroxynitrates in the southeastern United States and implications for the fate of NOx, Atmos. Chem. Phys., 15, 11257-11272, https://doi.org/10.5194/acp-15-11257-2015, 2015.

Xu, L., Pye, H. O. T., He, J., Chen, Y., Murphy, B. N., and Ng, N. L.: Experimental and model estimates of the contributions from biogenic monoterpenes and sesquiterpenes to secondary organic aerosol in the southeastern United States, Atmos. Chem. Phys., 18, 12613-12637, https://doi.org/10.5194/acp-18-12613-2018, 2018.

Yan, C., Nie, W., Äijälä, M., Rissanen, M. P., Canagaratna, M. R., Massoli, P., Junninen, H., Jokinen, T., Sarnela, N., Häme, S. A. K., Schobesberger, S., Canonaco, F., Yao, L., Prévôt, A. S. H., Petäjä, T., Kulmala, M., Sipilä, M., Worsnop, D. R., and Ehn, M.: Source characterization of highly oxidized multifunctional compounds in a boreal forest environment using positive matrix factorization, Atmos. Chem. Phys., 16, 12715-12731, 10.5194/acp-16-12715-2016, 2016.

Yáñez-Serrano, A. M., Nölscher, A. C., Bourtsoukidis, E., Gomes Alves, E., Ganzeveld, L., Bonn, B., Wolff, S., Sa, M., Yamasoe, M., Williams, J., Andreae, M. O., and Kesselmeier, J.: Monoterpene chemical speciation in a tropical rainforest:variation with season, height, and time of dayat the Amazon Tall Tower Observatory (ATTO), Atmos. Chem. Phys., 18, 3403-3418, 10.5194/acp-18-3403-2018, 2018.

Yuan, B., Koss, A. R., Warneke, C., Coggon, M., Sekimoto, K., and de Gouw, J. A.: Proton-Transfer-Reaction Mass Spectrometry: Applications in Atmospheric Sciences, Chem Rev, 117, 13187-13229, 2017.

Yucuis, R. A., Stanier, C. O., and Hornbuckle, K. C.: Cyclic siloxanes in air, including identification of high levels in Chicago and distinct diurnal variation, Chemosphere, 92, 905-910, https://doi.org/10.1016/j.chemosphere.2013.02.051, 2013.

Zhang, H., Yee, L. D., Lee, B. H., Curtis, M. P., Worton, D. R., Isaacman-VanWertz, G., Offenberg, J. H., Lewandowski, M., Kleindienst, T. E., Beaver, M. R., Holder, A. L., Lonneman, W. A., Docherty, K. S., Jaoui, M., Pye, H. O. T., Hu, W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Guo, H., Weber, R. J., de Gouw, J., Koss, A. R., Edgerton, E. S., Brune, W., Mohr, C., Lopez-Hilfiker, F. D., Lutz, A., Kreisberg, N. M., Spielman, S. R., Hering, S. V., Wilson, K. R., Thornton, J. A., and Goldstein, A. H.: Monoterpenes are the largest source of summertime organic aerosol in the southeastern United States, 115, 2038-2043, 10.1073/pnas.1717513115 %J Proceedings of the National Academy of Sciences, 2018.

Zhu, J., Penner, J. E., Yu, F., Sillman, S., Andreae, M. O., and Coe, H.: Decrease in radiative forcing by organic aerosol nucleation, climate, and land use change, Nature Communications, 10, 423, 10.1038/s41467-019-08407-7, 2019.

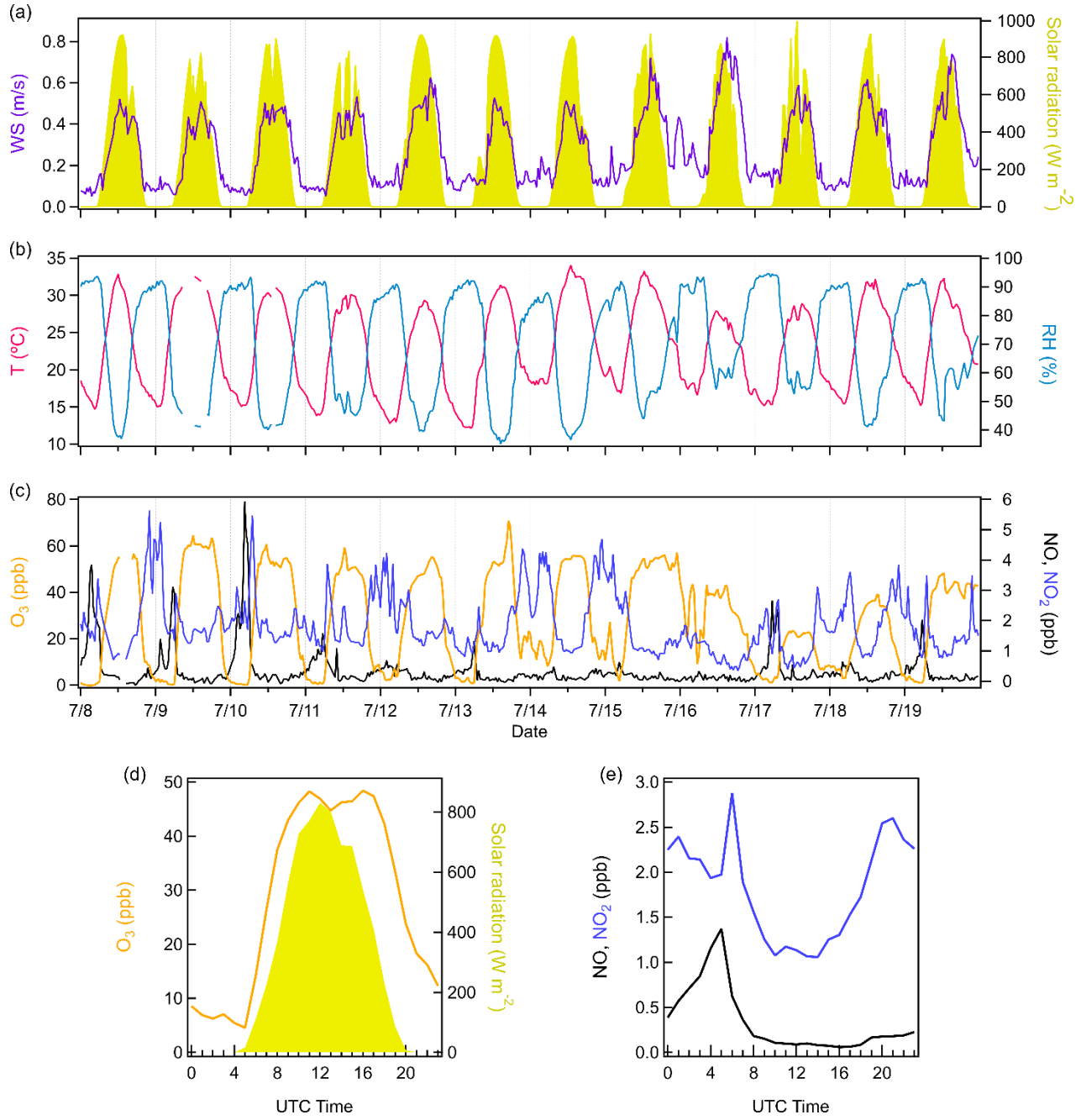


Figure 1. Variations of meteorological conditions and trace gases. (a) Time series of wind speed and solar radiation. (b) Time series of temperature and relative humidity. (c) Time series of O_3, NO, and NO_2. (d) Diurnal cycles of O_3 and solar radiation. (e) Diurnal cycles of NO and NO_2.

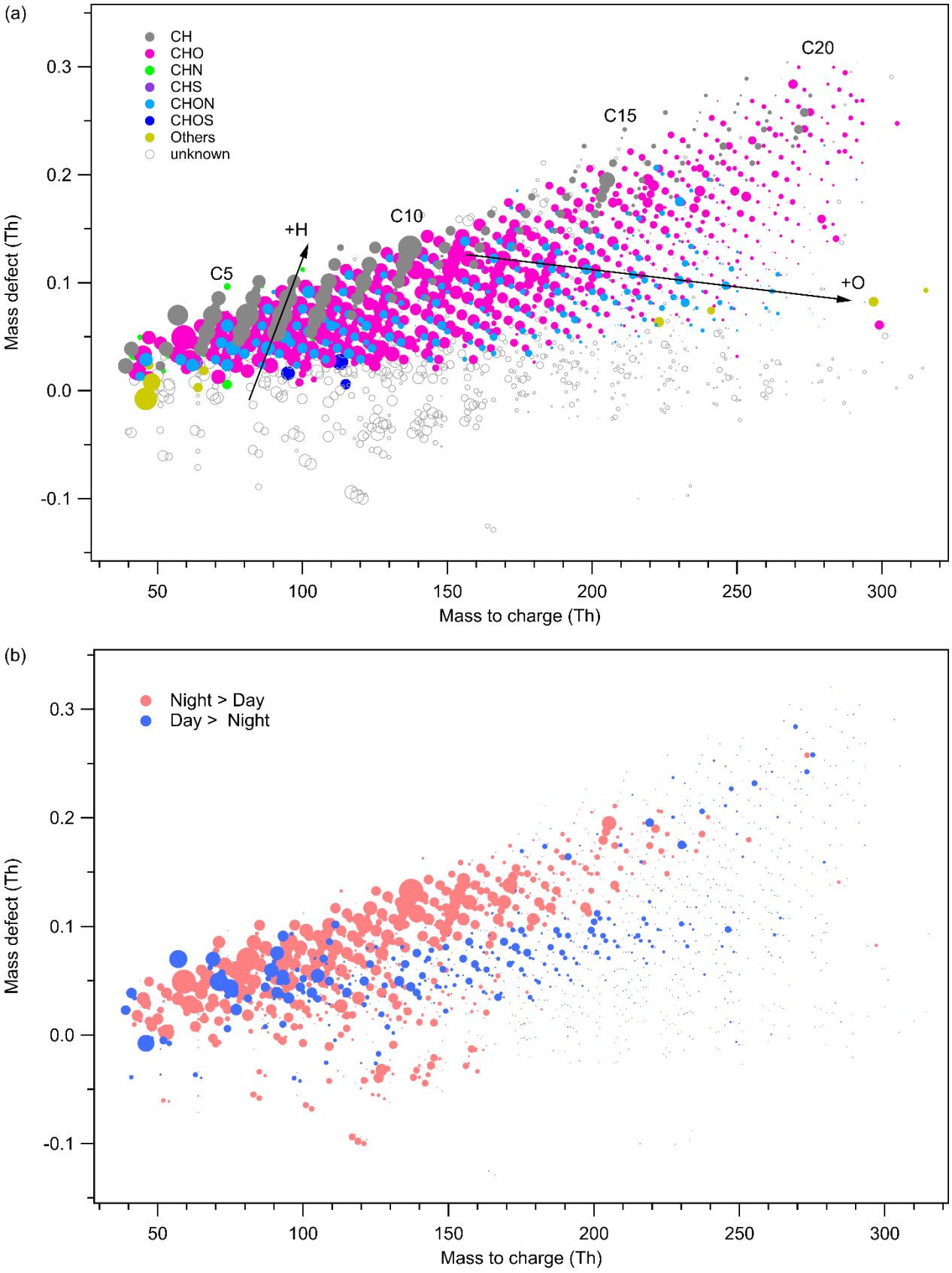


Figure 2. Mass defect plot of the ions identified by high-resolution analysis of Vocus PTR-TOF data set. The x-axis shows the mass to charge ratio and the y-axis shows the mass defect, which is the deviation of the exact mass from the nominal mass. Data points in (a) are color-coded by ion family (CH, CHO, CHN, CHS, CHON, CHOS) and sized by the logarithm of peak area. Data points in (b) are shown in pink when signals are higher during nighttime and in blue when daytime signal is higher. The size corresponds to the difference of daytime and nighttime signal for the molecule. It should be noted that ions < 35 Th are detected at a much-reduced efficiency due to a high-pass band filter in the BSQ.

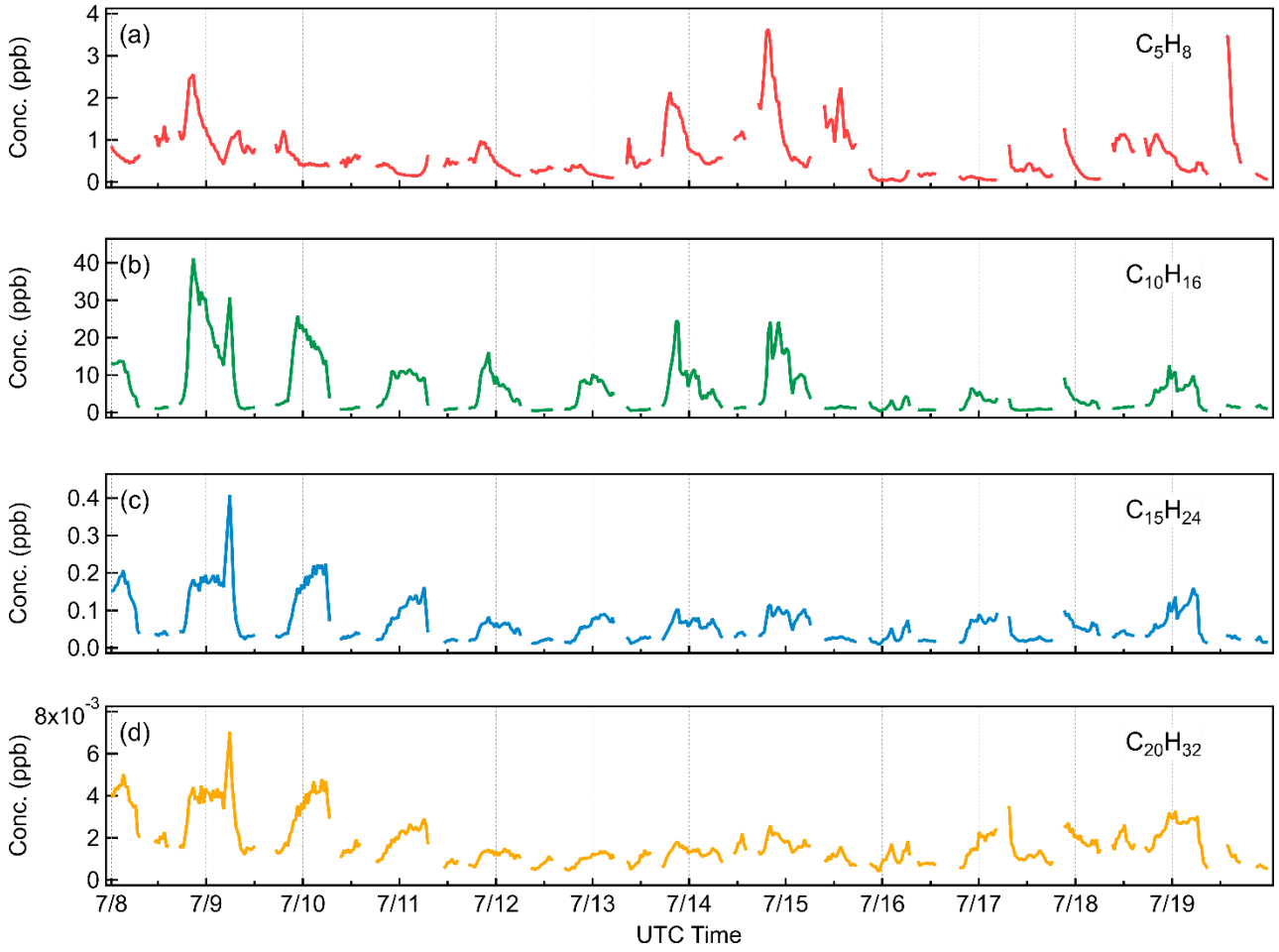
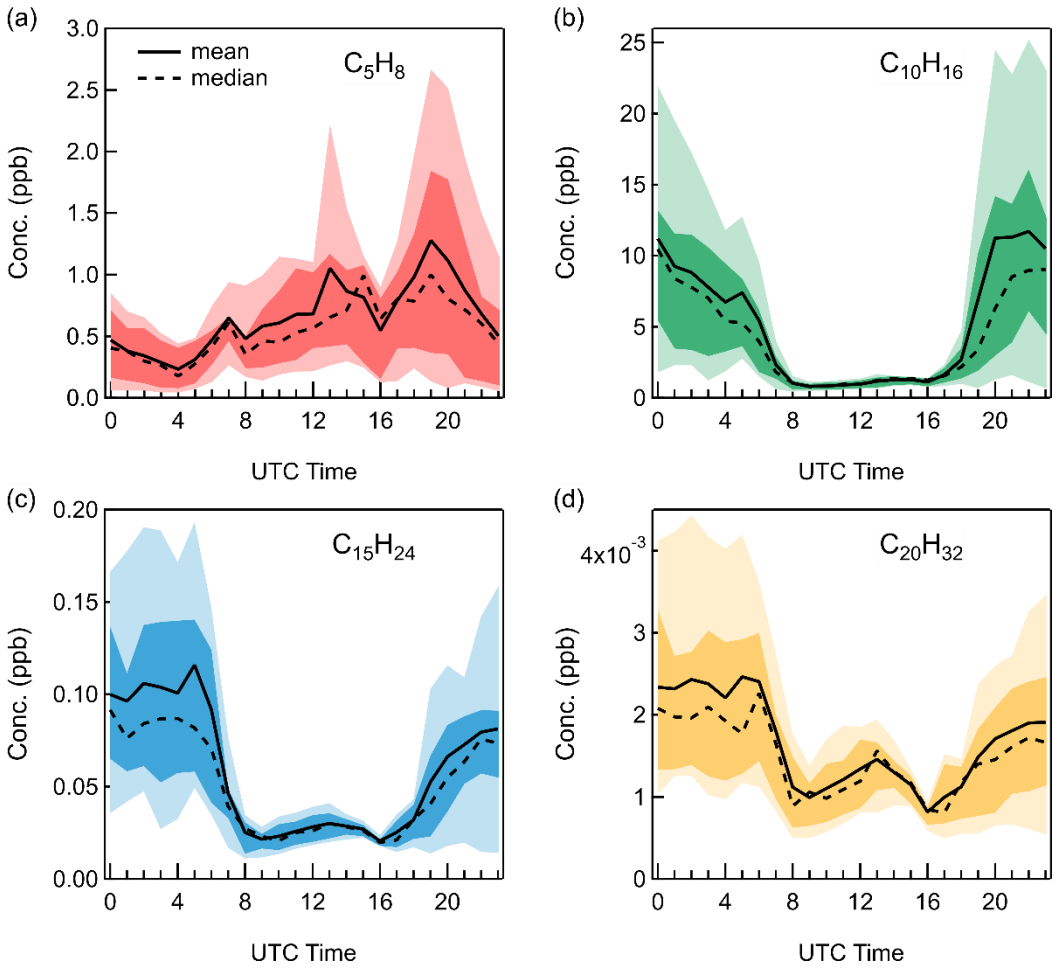


Figure 3. Time series of (a) C_5H_8, (b) $C_{10}H_{16}$, (c) $C_{15}H_{24}$, and (d) $C_{20}H_{32}$.



802

803 **Figure 4. Diurnal cycles of (a) C_5H_8, (b) $C_{10}H_{16}$, (c) $C_{15}H_{24}$, and (d) $C_{20}H_{32}$, with the 10th, 25th, 75th, and 90th percentiles**
804 **shown in the shaded area.**

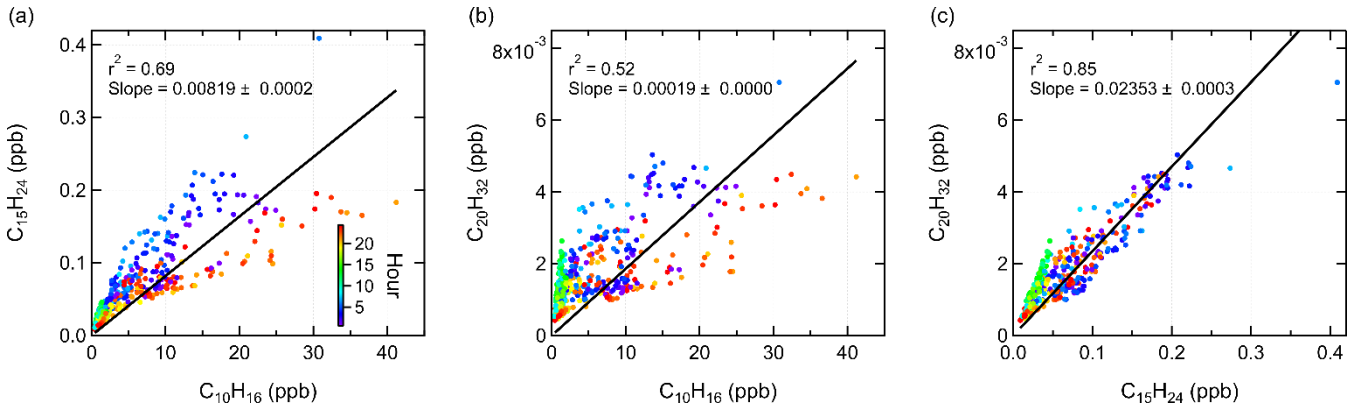


Figure 5. Scatter plots of (a) $C_{15}H_{24}$ vs. $C_{10}H_{16}$, (b) $C_{20}H_{32}$ vs. $C_{10}H_{16}$, and (c) $C_{20}H_{32}$ vs. $C_{15}H_{24}$, colored by time of the day.

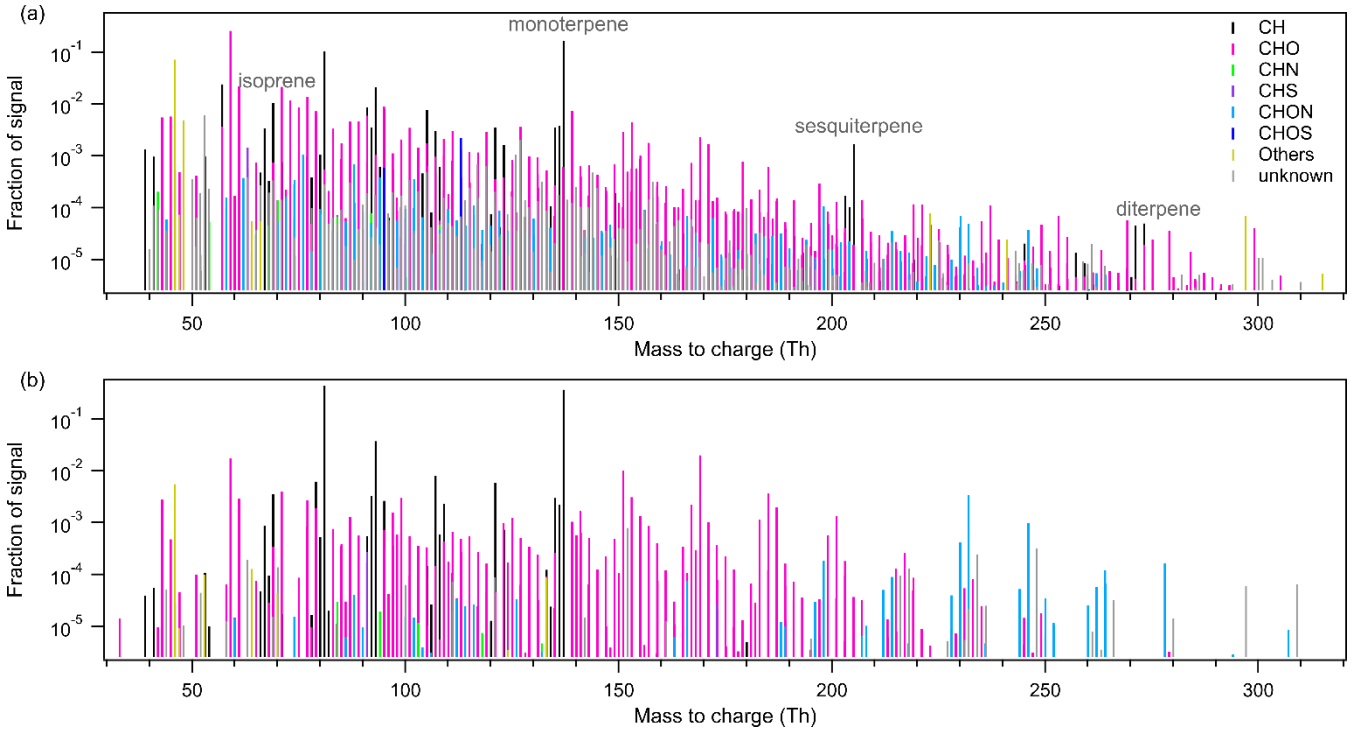
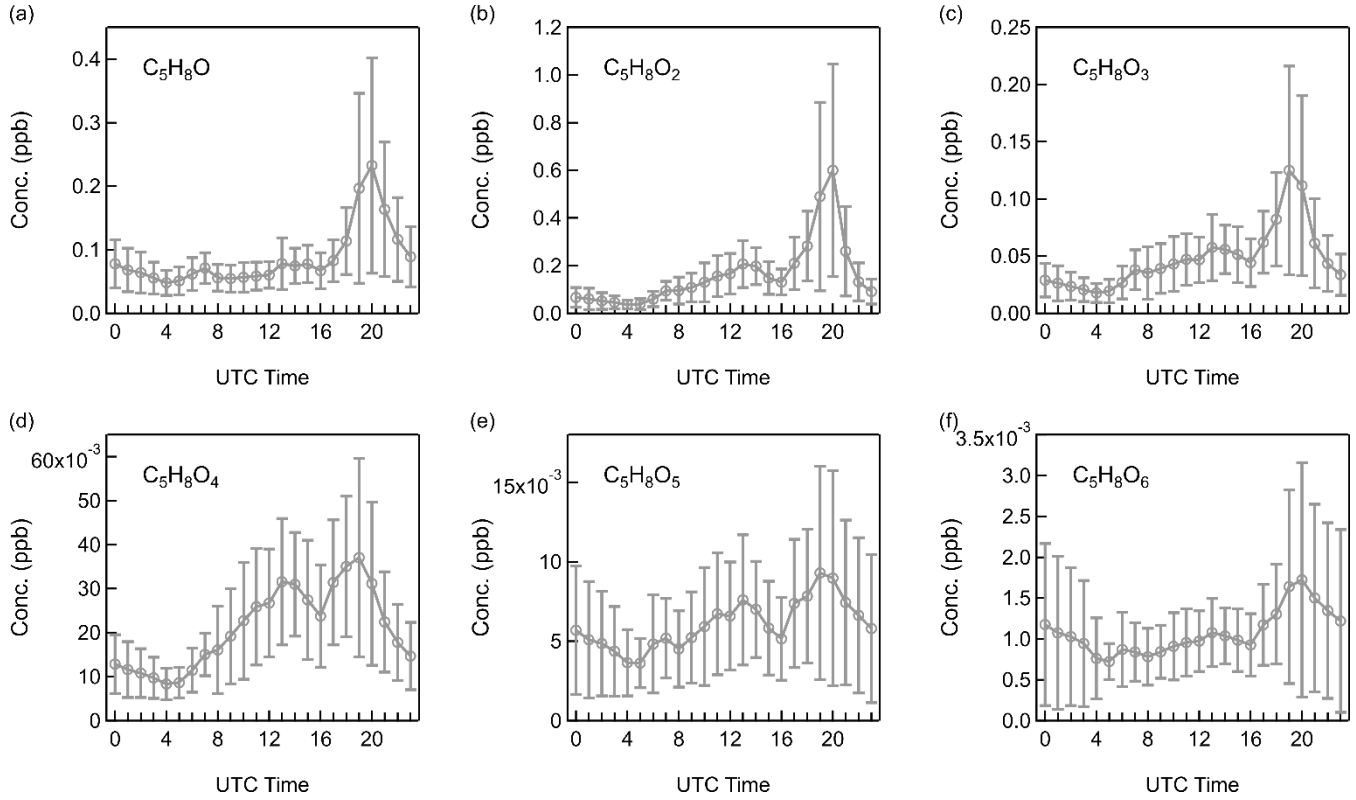


Figure 6. Comparison of ambient average high-resolution mass spectra with those from α-pinene oxidation experiments in the COALA chamber. (a) ambient observations in the Landes Forest; (b) α-pinene ozonolysis with NO_x.



811

812 **Figure 7. Diurnal patterns of non-nitrate isoprene oxidation products: (a) C_5H_8O, (b) $C_5H_8O_2$, (c) $C_5H_8O_3$, (d) $C_5H_8O_4$,**

813 **(e) $C_5H_8O_5$, and (f) $C_5H_8O_6$.**

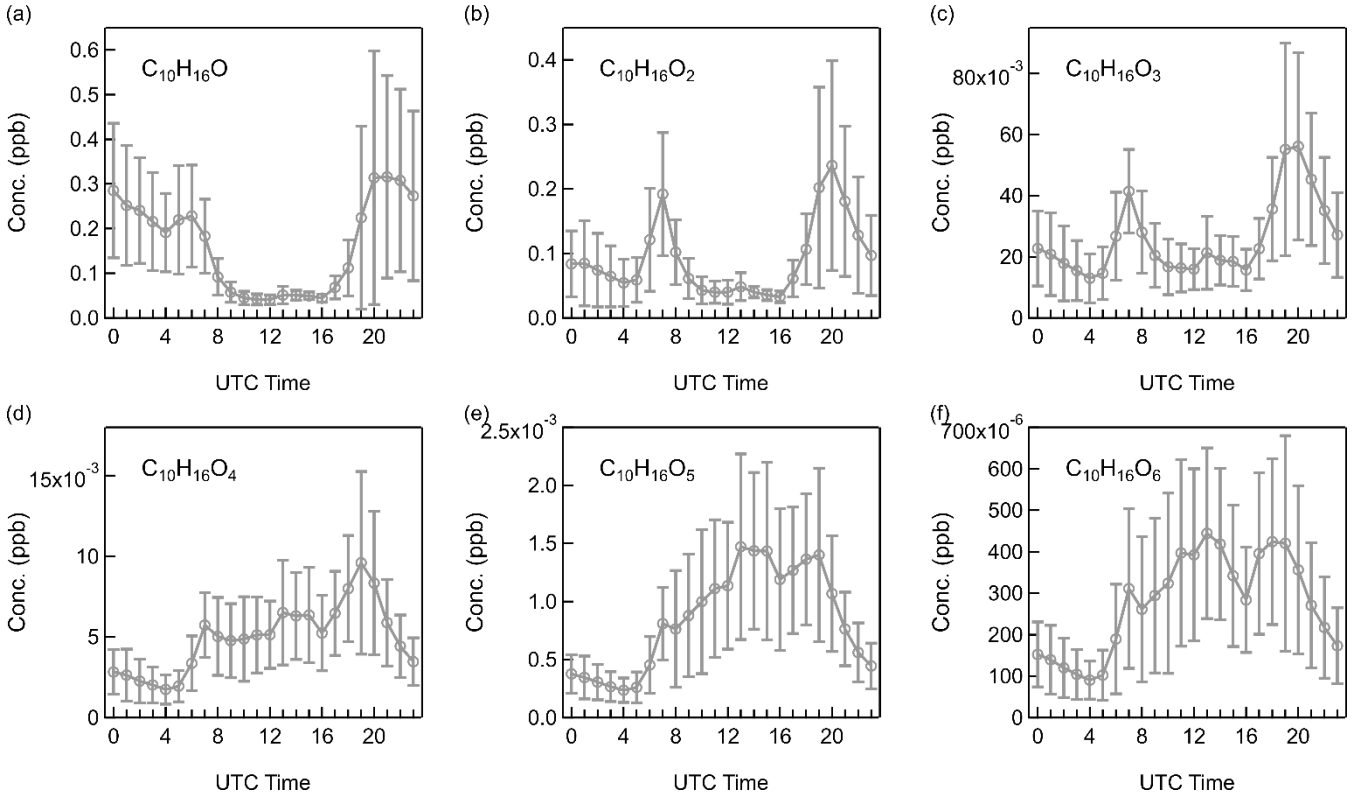
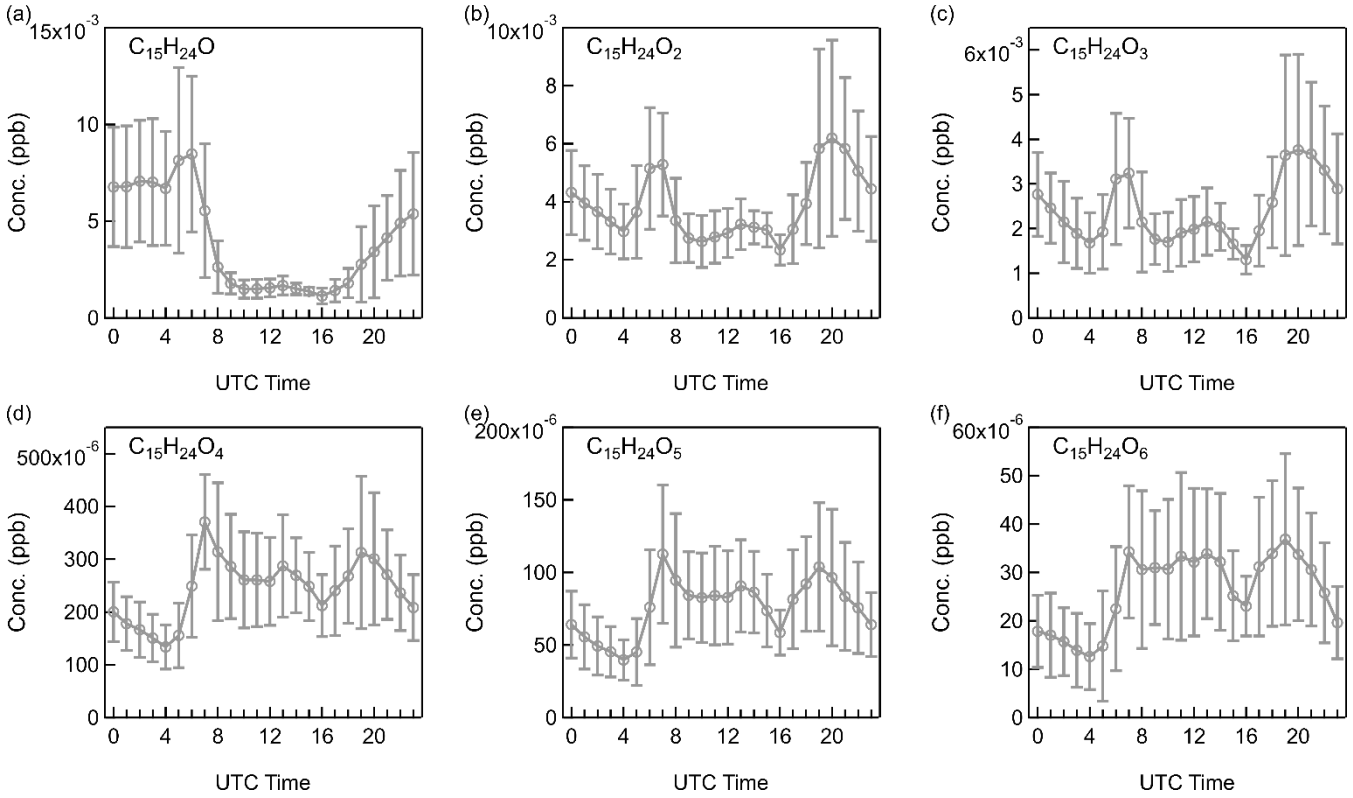
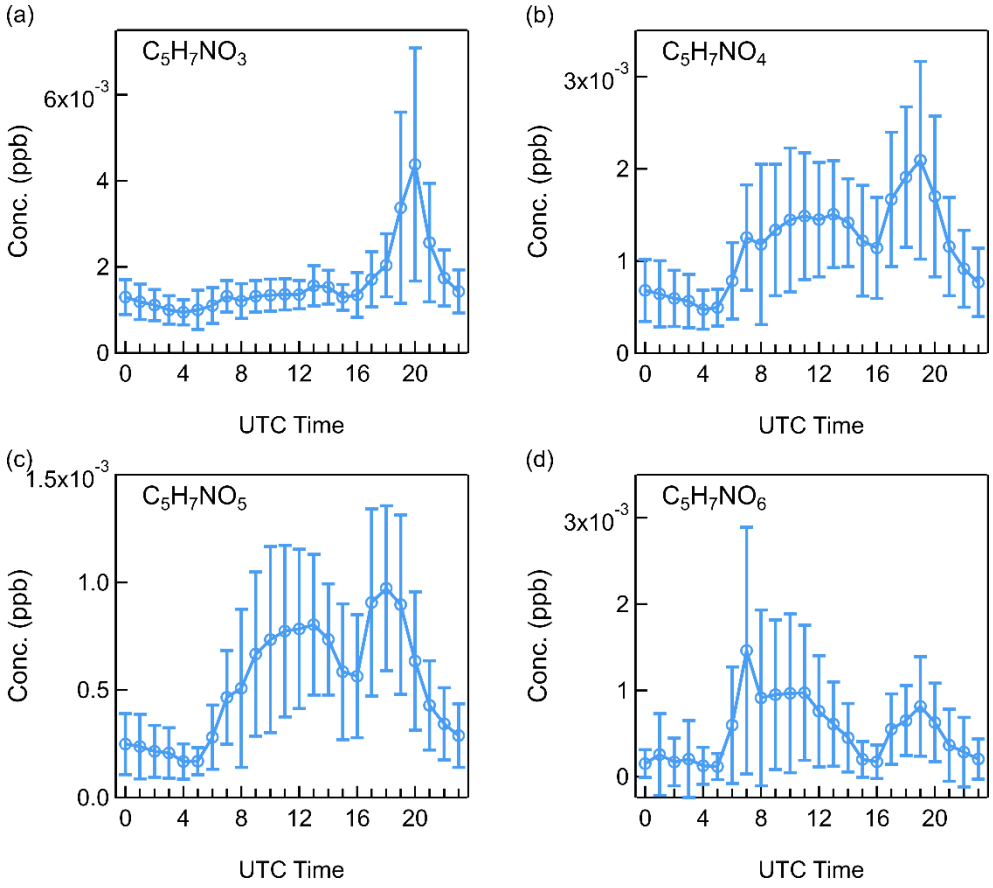


Figure 8. Diurnal patterns of non-nitrate monoterpene oxidation products: (a) $C_{10}H_{16}O$, (b) $C_{10}H_{16}O_2$, (c) $C_{10}H_{16}O_3$, (d) $C_{10}H_{16}O_4$, (e) $C_{10}H_{16}O_5$, and (f) $C_{10}H_{16}O_6$.



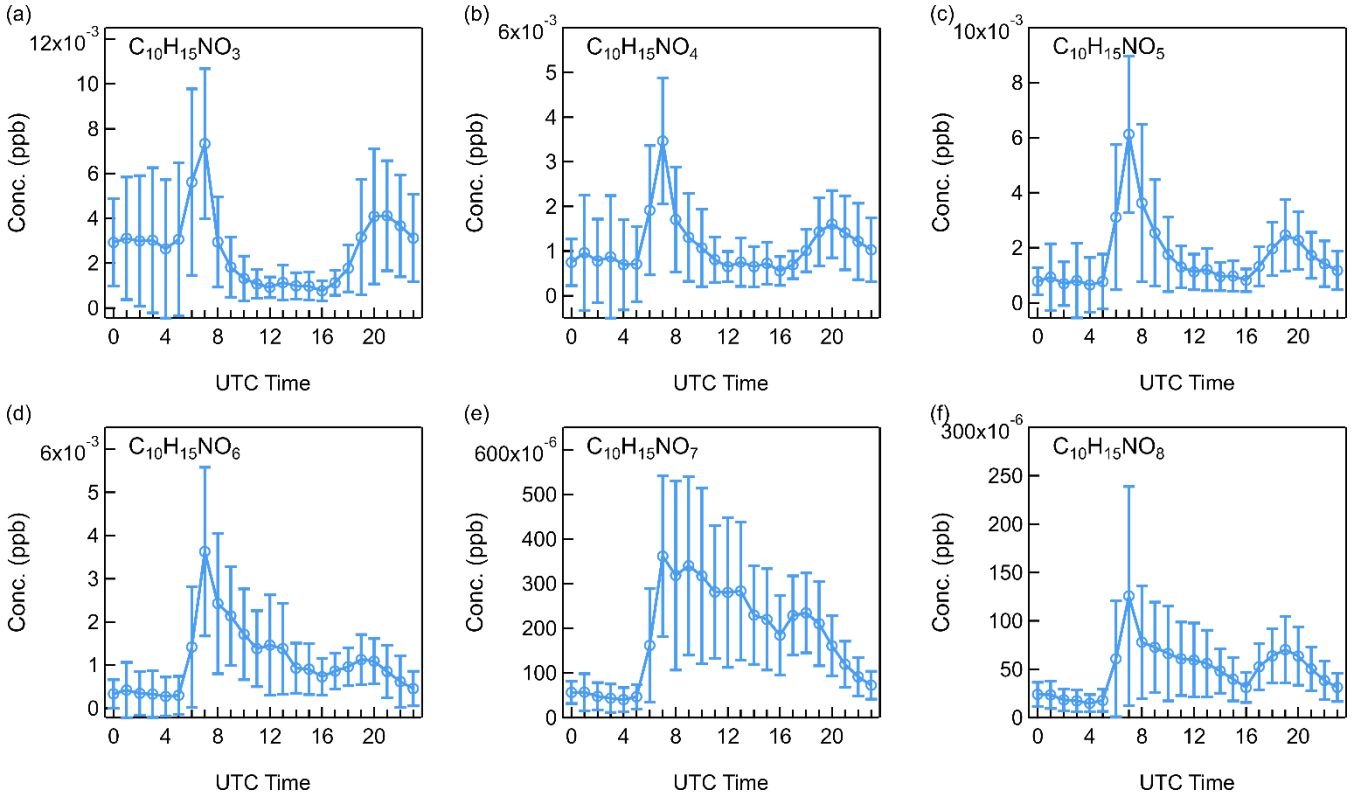
817

818 **Figure 9. Diurnal patterns of non-nitrate sesquiterpene oxidation products: (a) $C_{15}H_{24}O$, (b) $C_{15}H_{24}O_2$, (c) $C_{15}H_{24}O_3$, (d)**
819 **$C_{15}H_{24}O_4$, (e) $C_{15}H_{24}O_5$, and (f) $C_{15}H_{24}O_6$.**



820

821 **Figure 10. Diurnal patterns of isoprene-derived organic nitrates: (a) $C_5H_7NO_3$, (b) $C_5H_7NO_4$, (c) $C_5H_7NO_5$, and (d)**
822 **$C_5H_7NO_6$.**



823

824 **Figure 11. Diurnal patterns of monoterpene-derived organic nitrates: (a) $C_{10}H_{15}NO_3$, (b) $C_{10}H_{15}NO_4$, (c) $C_{10}H_{15}NO_5$, (d)**
825 **$C_{10}H_{15}NO_6$, (e) $C_{10}H_{15}NO_7$, and (f) $C_{10}H_{15}NO_8$.**