

*For clarity, referees' comments are written in black, our comments to each concern are written in blue and extractions from the original paper are written in green.*

## **Anonymous Referee #1**

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The manuscript by Villanueva et al. aims at contributing to the interpretation of the role of mineral dust in the formation of ice clouds. In particular, the authors want to demonstrate that the day-to-day the co-variability between mineral dust concentration and cloud glaciation can be used in future as a proxy to better understand heterogeneous freezing mechanism.

First of all I must admit that the manuscript is hard to read in several parts, and it takes more than one reading to make sure that the reported content is fully understood. The assessment of the most suitable datasets utilized for the study of heterogeneous freezing as well as the statistics related to the relationship between ice occurrence, updrafts and dust concentration are very interesting. Nevertheless I am not fully sure that this manuscript demonstrates the value of the day-to-day co-variability as a proxy of the effect on cloud glaciation.

We thank Referee 1 for thoroughly reading the manuscript and for his/her detailed comments. We believe that the manuscript is now easier to read, after the changes we made in response to the Referee's comments. We also thank the reviewer for his/her encouraging comments about the statistics related to the ice occurrence and dust concentrations, which are indeed the main new idea that we want to present in the manuscript. We regret that the referee is still not convinced of the day-to-day statistical approach, but we hope that our response may satisfy the reviewer's concerns.

Below I report my general comments.

- First of all, the authors have selected data from many difference sources, MACC and ERA interim meteorological reanalysis CALIPSO-GOCCP, DARDAR products from the A-Train satellite constellation. What are the effects on the final results presented in the manuscript of combining these datasource with different speculation (resolutions, sampling, uncertainties, ..)?

Concerning the different product resolutions, the choice of 3 K bins is based mainly on the resolution of the CPR (480 m) — used in the DARDAR product — and the original 3 K bins of the CALIPSO-GOCCP product.

In our methodology, a higher resolution (e.g., 1 K bins) would result in a more heterogeneous distribution of the sample size — for any given temperature and day, a larger fraction of gridboxes would contain no clouds — and therefore the statistical uncertainty in the day-to-day statistics would increase. Although the instantaneous satellite products could be resampled to 1 K bins, for the reanalysis the vertical levels would need to be interpolated. Furthermore, some additional technical disadvantages would be associated with a higher resolution. In general, increasing the resolution of the dataset combination would require a new methodology including interpolation of some products. To test if smaller bins could increase the confidence in the results by better constraining the temperature effect on cloud phase, we have performed a quick sensitivity study using the whole range 2007-2010 of the DARDAR-MASK product. In this quick study, we used 3 K bins and 1 K bins to quantify how the vertical resolution affects the variance of the ice-to-liquid ratio. This sensitivity study shows that at -30°C the temporal standard deviation of the FPR decreases by less than 10% in the 1 K bins, despite a decrease of almost 50% in the standard deviation of the temperature inside the temperature bin. Although this represents only a small subset of the data, it shows that a higher vertical resolution not necessarily leads to a decrease in the uncertainty.

Table R1. Temporal standard deviation (STD) of the ice to liquid ratio from the DARDAR-MASK product by different resolutions (RES) for the temperature bins. The daily values for the whole period 2007-2010 were used. Only global means are shown (96 lat×12 lon)

INTERVAL	RES	STD	VARIABLE
-30°C TO -33°C	3K	0.17	FPR_DARDAR
-30°C TO -31°C	1K	0.15	FPR_DARDAR
-31°C TO -32°C	1K	0.16	FPR_DARDAR
-30°C TO -33°C	3K	0.047	Temperature
-30°C TO -31°C	1K	0.028	Temperature
-31°C TO -32°C	1K	0.028	Temperature

Analogously, using a coarser vertical resolution (e.g., 6 K bins) would allow larger temperature variations. For example, a decrease of 3 K is roughly equivalent to a fivefold increase in INP concentrations (e.g., Niemand et al., 2012). However, because the typical range of the day-to-day variations of dust mixing-ratio is also large — about 1 order of magnitude — we expect that the variability of dust loading should dominate over temperature variations, given a constraint of at least 3 K.

The statistical distribution of the ice ratio also limits the resolution options. Single pixels values of cloud phase from the DARDAR and GOCCP products are binary (1 or 0). Therefore, a minimal sample size is required for the averaged cloud phase —within a certain temperature range, gridbox and percentile of dust— to achieve a normal distribution, which allows interpreting the correlation with dust loading directly.

We have now included this matter in the discussion section (without the table):

+(new Lines 517-528) ”Concerning the vertical resolutions of the different products, the choice of 3 K bins is based on the resolution of the CPR (480 m) — used in the DARDAR-MASK product — and the original 3 K bins of the CALIPSO-GOCCP product. Using a coarser vertical resolution (e.g., 6 K bins) would hinder the assessment of the role of dust as INP. For example, a decrease of 3 K in temperature is roughly equivalent to a fivefold increase in INP concentrations (e.g., Niemand et al., 2012). Because at the mid- and high-latitudes the typical standard deviation of the day-to-day dust mixing-ratio corresponds to roughly a fourfold increase from the mean (See supplement figure S.5), we expect that the variability of dust loading should dominate over temperature variations, given a temperature constraint of 3 K or less. The statistical distribution of the phase ratio also limits the resolution options. The cloud phase values for single pixels in the DARDAR and GOCCP products are binary (1 or 0). Therefore, a minimal sample size is required for the averaged cloud phase ratio — within a certain temperature range, gridbox and percentile of dust — to achieve a normal distribution along time and space, which allows interpreting the correlation with dust loading directly. For this reason, temperature bins smaller than 3 K result in a less normally-distributed cloud phase ratio.“

- The authors often do assumptions and simplifications (e.g., use of night time measurements only, neglecting of ice in the mixed phased clouds, . . .) which can strongly increase the uncertainty of the final results and limit the value of the data interpretation. Cloud phase is mainly regulated by temperature and so it is not clear to me why the authors considered only nighttime measurements. What’s the effect of this data selection on the final results?

As mentioned in section 3.1, we exclude the daytime retrievals from CALIOP to avoid the influence of noise from sunlight scattering on the retrievals (Li et al., 2017). This is a known issue present in the daytime CALIOP retrievals and this data selection is a necessary step to prepare a consistent dataset (e.g., due to sunlight scattering day and night-time CALIOP retrievals are not directly comparable for our purposes). Nevertheless, the first experiments with the original data showed that the inclusion of daytime retrievals lead to lower FPR-dust correlations (not shown).

As for the neglectation of ice in mixed-phase clouds, this was only done in the FPR\_DARDAR\_ALT variable to study the differences between the CALIPSO-GOCCP and DARDAR products. The final results using the GOCCP product are not affected by this simplification.

- The scope of the manuscript is to demonstrate that the day-to-day co-variability of dust concentration ice cloud glaciation may be used to quantify the role of dust aerosol on the cloud thermodynamic phase around the globe. I understand that, in order to use the day-to-day variability, the authors removed the seasonal component subtracting the monthly means. I am not sure this is sufficient to remove all the possible variabilities which can affect the selected data. Weather variability for example occurs on a scale of 5-6 days. Can the authors explain how they can assure the data are not affected by any other relevant variability cycle?

As correctly pointed out by the reviewer, the day-to-day variability may still include the effects of weather variability. To disentangle the weather variability from the day-to-day variability is indeed an interesting idea. Nevertheless, the number of data samples available is not enough to additionally subclassify the retrievals according to local weather states and, at the same time, to also assess the effect of day-to-day dust variability.

We are convinced that it is still possible to study the relationship between dust and cloud phase even without excluding cycles shorter than one month. Although cloud phase may be affected by such cycles (e.g., more liquid clouds at convective fronts and more cirrus clouds at the detrainment regions), it is still possible to distinguish between dusty and non-dusty conditions at each point of the weather cycle. Therefore, once we average over the weather cycle — using monthly means inside each dust percentile — it is still possible to observe the dust-cloud-phase relationship.

These points have been added to the discussion section:

+(new Lines 529-535): “As mentioned in section 3.5, we excluded the seasonal component of the dust-cloud-phase correlation by calculating the deciles independently for each month of the year. However, shorter cycles (e.g., weather variability) may still have an influence in the variability of dust and cloud phase. Although the cloud phase may be affected by such cycles (e.g., more liquid clouds at convective fronts and more cirrus clouds at the detrainment regions), it is still possible to distinguish between dusty and non-dusty conditions at each point of the weather cycle. Therefore, once we average over the weather cycle — using monthly means inside each dust percentile — we expect the dust-cloud-phase relationship to be dominated by the microphysical effect of dust on cloud phase.”

- In addition, it is not clear to me from the reported description whether the 3K binning can smooth the day-to-day variability (though the binning was needed with respect to the considered dataset)

The 3 K can be understood as averaging the satellite pixels between the height of two different isotherms. These isotherms change on a day-to-day basis as estimated by the reanalysis. Therefore, even small day-to-day changes in temperature (e.g., 1 K) are traduced in slight changes in the isotherm heights and in the interval for which the 3 K bins are calculated. Therefore, we do not expect the effect of day-to-day changes in temperature to be an important source of uncertainty in our analysis.

- ... or anyhow mix different observation scenarios, i.e. high dust content and low content but at the same temperature.

This is an interesting concern. However, the 3K binning is unlikely to produce a significant mixing between different dust scenarios. Even though aerosol plumes can absorb short and longwave radiation, local differences are usually less than 3 K. Even Saharan dust layers has been found to produce temperature differences of only about +2 K in the North Atlantic (Wang and Liu, 2014).

Although cases of high dust concentration over land could produce higher temperature differences — and therefore a temperature inversion —, we do not expect such rare cases to introduce a significant systematic bias in the results. Neither do we expect temperature inversions to play a significant role at temperatures below -15°C, except perhaps near the poles.

- I think the description in section 3.2 must be clearer.

—“3.2 Regridding and rebinning: Temperature levels and  $1.875^\circ \times 30^\circ$  gridboxes

*It will become clear in Sect. 4.2. That the cloud thermodynamic phase is mainly a function of temperature. Anticipating this, temperature bins of 3 K each were used as vertical coordinate throughout the study. The temperature profiles were obtained from the ECMWF-AUX reanalysis for the DARDAR and CLDCLASS products and from the MERRA reanalysis for the GOCCP product.*

*To fill the horizontal gaps between the satellite orbits, we regridded the dataset into a Gaussian T63 grid, aggregating 16 gridboxes along the longitude ( $1.875^\circ \times 30^\circ$ ; lat $\times$ lon). The Gaussian T63 grid is commonly used in Global Climate Models (Randall et al., 2007) and facilitates future comparisons with global simulations of cloud thermodynamic phase. In section 4.4 and onwards, latitude bands of  $30^\circ$  are used to allow a direct comparison with previous studies (Zhang et al., 2018).“*

We have changed the description to:

+(new Lines 178-190)“3.2 Regridding and rebinning: Temperature levels and  $1.875^\circ \times 30^\circ$  gridboxes

*It will become clear in Sect. 4.2. that the cloud thermodynamic phase is mainly a function of temperature. Anticipating this, temperature bins of 3 K each were used as a vertical coordinate throughout the study. The temperature profiles were obtained from the ECMWF-AUX reanalysis for the DARDAR and CLDCLASS products and from the MERRA reanalysis for the GOCCP product. Thus, the same temperature information was used for each product algorithm and for its postprocessing. The temperature at each profile pixel is interpolated using the information from the reanalyses and between  $-42^\circ\text{C}$  and  $+3^\circ\text{C}$ , the pixels are averaged into 3 K intervals.*

*For each of these 3 K intervals, the pixels are then averaged horizontally in a process known as regridding. We regridded the dataset into a Gaussian T63 grid, aggregating 16 gridboxes along the longitude ( $1.875^\circ \times 30^\circ$ ; lat $\times$ lon) to better fill the horizontal gaps between the satellite orbits. The Gaussian T63 grid is commonly used in Global Climate Models (Randall et al., 2007) and facilitates comparisons with global simulations of the cloud thermodynamic phase. In section 4.4 and onwards, latitude bands of  $30^\circ$  are used to allow a direct comparison with previous studies (Zhang et al., 2018).“*

- Many time the authors state that the presented results can be affected by assumptions or effects not considered but also that several properties, which at regional scale may have significant difference, should be reconciled by the fact that the static stability is calculated at the global scale. This is not true for all the variables, RH is an example. There might be strong variations of RH at regional scale in one hemisphere only, which can affect the value of the interpretation of results. Sampling uncertainties are mentioned a few times by the authors themselves, though these are never quantified.

The constraints on statistic stability and RH were intended to constrain the temporal and not the regional variability. Therefore, intervals used as constraints are based on the day-to-day meteorological variations and do not consider the regional variability. We agree with the referee that the uncertainty from the regional variability may be a relevant source of uncertainty. However, such regional variations are at least partially accounted by the standard deviation (zonal) shown in Figures 8-11 (error bars). The focus of the study is not the local regional differences but rather the potential hemispheric differences. Therefore, a deeper investigation of smaller-scale regional correlations is out of the scope of this study.

- For example, if the regions where the data sample is larger and more complete is resampled to reduce the amount and obtain a dataset of homogenous size across the zonal regions what the effect would be?

The suggestion is to change the gridbox size (“resample”) at each latitude to obtain the same number of data point inside each gridbox. The result would be a highly irregular grid, however, as the reviewer suggests, the data distribution should be much more homogeneous between such gridboxes. Indeed, there would

be some statistical advantages of this approach (e.g., the zonal averaging would be associated with a lower variance). Nevertheless, new problems would arise, as for example a bias towards regions with higher cloud cover — and therefore a larger sample size. Moreover, our analysis is based on the assumption of a regular grid and we consider the implementation of such a challenging dynamic grid structure out of the scope of our study.

Below I report some detailed comments sometimes still of general breath.

- Line 177-180: Re-gridding operated here generates also a degradation of the horizontal resolution whose effect on the provided analysis is not quantified. It is not clear to me if this is a real advantage or not.

—(Lines 183-186) *“To fill the horizontal gaps between the satellite orbits, we regridded the dataset into a Gaussian T63 grid, aggregating 16 gridboxes along the longitude ( $1.875^\circ \times 30^\circ$ ; lat $\times$ lon). The Gaussian T63 grid is commonly used in Global Climate Models (Randall et al., 2007) and facilitates future comparisons with global simulations of cloud thermodynamic phase. In section 4.4 and onwards, latitude bands of  $30^\circ$  are used to allow a direct comparison with previous studies (Zhang et al., 2018). “*

- As stated in the text, the purpose of this regridding is to ensure enough datapoints in each gridbox. It is correct that this degrades the horizontal resolution. However, the analysis is focused on the latitudinal differences — like the North-South hemispheric contrast — rather than on regional differences.

In general, due to the explorative nature of the approach, we did not pursue a detailed sensitivity study of the dataset configuration (e.g., temporal, vertical and horizontal resolution). However, during the initial screening, higher horizontal resolutions (e.g.  $2 \times 2^\circ$ ) resulted in a too-large zonal variance, due to the binary-nature of the ice-to-liquid ratio. In other words, gridboxes with only few retrievals per day were strongly biased towards single clouds —liquid(0) or ice(1)—, so that the zonal standard deviation was higher than the changes due to dust loading variability.

- Line 185: replace “in the study” with the “in the study by Huang et al”
  - Replaced (Line 194)
- Line2 190-192: please clarify that the choice to ignore ice in mixed phase clouds is an advantage according to the approach you are adopting but also that the authors cannot be sure this has not an impact on the final results.

—(Lines 196–198) *“In this alternative definition, which we will call ALT- DARDAR, only gridboxes ( $1.875^\circ \times 30^\circ \times 3$  K) filled with ice pixels are considered as ice (fully glaciated), so that just a single liquid pixel is enough to define a gridbox as liquid (not fully glaciated). One advantage of this marginal definition is that it ignores cloud ice in mixed-phase clouds, which is mostly only detected as such by the DARDAR-MASK product and neglected by the CALIPSO-GOCCP product. “*

We have now clarified this statement by adding the following:

+(new Lines 203-204) *“However, this neglecton of ice in mixed-phase clouds is only carried out to clarify the differences between the products.”*

- Lines 264-266: the authors limit their investigation to the altocumulus clouds: can they quantify the impact of this choice on the final results?

—(Lines 261-266) *“Fig. 4a-b show for the same segment the cloud volume cover (CALIPSO-GOCCP) of clouds classified (2B-CLDCLASS) as cirrus or altocumulus (Fig. 4a) and as altostratus or stratocumulus (Fig. 4b). These cloud types are frequently thin enough to be penetrated by lidar and radar systems and are therefore a good target to study cloud glaciation processes (Bühl et al., 2016; D.Zhang et al., 2010b). “*



- We are a bit confused by this remark. We are afraid we didn't clarify this point enough. We do not limit the analysis to altocumulus only. Furthermore, these lines refer only to the case study. We have added an introductory statement to improve clarity.

+(new Lines 273-280) *"To better understand the differences between the ice-to-liquid ratio retrieved in the DARDAR-MASK and CALIPSO-GOCCP product, this section provides a detailed case study of a stratiform cloud scenario, in which four stratiform cloud types from the CloudSat classification are included — stratocumulus (low-level clouds), altostratus and altocumulus (mid-level clouds), and cirrus (high-level clouds). Although not present in the case study, Nimbostratus are included in the analysis of cloud phase as well and are particularly important in the high latitudes. Stratus clouds are defined for temperatures above 0°C; therefore, they are not relevant for this study. Finally, the horizontal extension of Cumulus and Deep Convection clouds is very low compared to the stratiform clouds and can be therefore ignored in our study, especially outside the tropics (Sassen and Wang, 2008)."*

- Line 322: I think this simplification can create confusion only.

*—"Additionally, in the DARDAR algorithm water can be still classified as ice at +1.5°C due to the melting layer being set to a wet-bulb temperature ( $T_w$ ) of 0°C. This allows the detection of ice at temperatures slightly above 0°C dry-bulb temperatures (named simply temperature in this work)."*

We understand this concern, which was already pointed out in the quick reports — "Line 315: please use  $T_w$ , remove the sentence in brackets and do not adopt the proposed simplification throughout the text". We agree that the role of  $T_w$  (wet-bulb) and  $T_d$  (wet-dry) in the DARDAR-MASK product is somewhat confusing. However, the use of  $T_w$  throughout the analysis of the DARDAR product would aggravate the problem.  $T_w$  is only used in the classification algorithm of the DARDAR-MASK product, and its only purpose here is to define the height below which ice is not acceptable anymore. In other words,  $T_w$  has only an influence in the  $T_d$  range +3°C to 0°C, where it can extend the temperature interval where ice is allowed in the classification. This temperature bin is only briefly mentioned to explain the high temperature end in Figure. 5 and is not further considered afterwards."

- Lines 323-325: though concentration of dust is lower at high altitudes, this does not necessarily indicates that this is due to lower temperatures; this sentence create confusion and solve in a few words a more complicated issue which involves also many other factors, such as atmospheric dynamics and radiative budget. For example, there might be a feedback mechanism influencing the top altitude of aerosols. I think the sentence must be rephrased or otherwise removed.

*— (Lines 323-325) "Additionally, the average fine-mode dust mixing-ratio is also shown in Fig. 5. At 0°C the mixing-ratio is five times higher than at -42°C (note the logarithmic right y-axis). This reflects the fact that dust mixing-ratios tend to be lower at higher altitudes where temperatures are lower. However, there are important exceptions to this, such as in the long-range transport of dust layers over the ocean."*

*We have rephrased as follows:*

+(new Lines 344-348) *" Additionally, the average fine-mode dust mixing-ratio is also shown in Fig. 5. At the height of the 0°C isotherm, the mixing-ratio is on average higher than at the -42°C isotherm (note the logarithmic right y-axis). This reflects the fact that, on average, dust mixing-ratios tend to be higher near the dust sources at the surface. However, this does not imply any general relationship between dust and temperature. Moreover, instant vertical profiles of dust loading and temperature may differ greatly from this average, especially in the long-range transport of dust plumes."*

- Line 330: the authors should clarify the reason for the supposed correlation between the maxima observed in the NH and at the tropics, and how the transport of ice clouds downward may occur. Can this be related to any wave activity at the synoptic scale?

*—"These maxima are probably associated with the enhanced homogeneous freezing in the tropics at temperatures below -40°C and the resulting downward transport of cloud ice."*

First of all, we must clarify that is not the scope of the paper to clarify the mechanisms of ice production in the tropics. Our assumption is based on previous global climate simulations studies, where the main source of cloud ice below the  $-42^{\circ}\text{C}$  isotherm is the ice detrained from convective outflows (Gasparini and Lohmann, 2016). These ice particles may be produced by the rapid injection of cloud droplets to temperatures lower than  $-42^{\circ}\text{C}$ . Ice particles tend to grow and sediment faster than cloud droplets, and the associated downdrafts may enhance the downward transport of the detrained cloud ice. Because convection in the mid-latitudes is closely associated with cyclone activities, the synoptic-scale may indeed play an important role in the large-scale correlation between updrafts and cloud ice. We encourage future studies seeking to investigate this possibility.

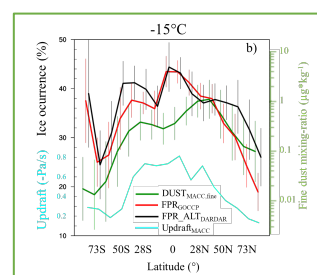
- Lines 342-343: I do not see the steep increase at the Northern Pole. I ask the authors to clarify.

We are a bit puzzled by this remark. The steep increase described in the text is refers to the Southern Pole. We have clarified this as following.

—( Lines 348-349)“For both variables, a local minimum near  $73^{\circ}\text{S}$  is followed by a steep increase at  $84^{\circ}\text{S}$ ”

This was changed to

+(new Lines 363-365) ”Moreover, the FPRGOCCP at  $-15^{\circ}\text{C}$  is lower than the FPR\_ALTDARDAR at the southern mid-latitudes and northern high-latitudes. In the southern high latitudes, for both variables, a local minimum near  $73^{\circ}\text{S}$  is followed by a steep increase at  $84^{\circ}\text{S}$ .”



- Lines 345-347: I think the authors may remove this lines, too conjectural; digressions are not needed in that part of the manuscript.

(Lines 351-353)—“The predominance of ice clouds in Antarctica has been already pointed out earlier in the literature (Ardon-Dryer et al., 2011; Bromwich et al., 2012). Incoming air masses from the ocean may carry higher concentrations of INP like biogenic aerosol (Saxena, 1983), Patagonian soil dust or Australian black carbon (Bromwich et al., 2012). “

We have removed these lines.

- Lines 348-350: how much does the number of data influence your conclusions at the South Pole? The authors should discuss this aspect in the paper.

(Lines 355-358)—“Similarly, it has been shown that the orographic forcing in Antarctica can lead to high ice water contents for maritime air intrusions (Scott and Lubin, 2016). In other words, maritime air intrusions associated with higher temperatures, higher concentrations of INP and stronger vertical motions could explain the observed pattern in the southern polar regions.”

We recognize that these lines are somewhat speculative. We agree with the referee, that the low sample size near the South Pole (Fig. 3 and supplement material s14.b) together with the low altitude of the  $-15^{\circ}\text{C}$  isotherm (s12.b) hinders more robust statistics. For example, at  $-15^{\circ}\text{C}$ , the zonal standard deviation of the FPR significantly increases from  $60^{\circ}\text{S}$  towards the South Pole — from about  $\pm 0.08$  to  $\pm 0.16$  in Fig.6a — at the same time that the sample size decreases from 2200 to 300 (Fig.3).

We have now included this issue in the text.

+(new Lines 370-374) “However, the low sample size near the South Pole (**Fig. 3** and supplement material **S.14.b**) and the low altitude of the  $-15^{\circ}\text{C}$  isotherm (**S.12.b**) result in a lower confidence in the results for this region. For example, at  $-15^{\circ}\text{C}$ , the zonal standard deviation of the FPR significantly increases from  $60^{\circ}\text{S}$  towards the South Pole — from about  $\pm 0.08$  to  $\pm 0.16$  in Fig.6a — at the same time that the sample size decreases from 2200 to 300 (**Fig.3**). ”

- Lines 352 - 355: the correlation mentioned here between the updraft and the FPR looks not so strong, can the authors provide numbers (i.e. regression coefficient or any other statistical tests)?

—“The pattern of the mean large-scale vertical velocity (MACC reanalysis) of the clouds studied is particularly similar to the FPR at  $-15^{\circ}\text{C}$ . Moreover, the spatial correlation between large-scale updraft velocity at 500 hPa is positively correlated to the occurrence-frequency of ice clouds at  $-20^{\circ}\text{C}$  (Li et al., 2017a). In other words, both the dust mixing-ratio and the large-scale vertical velocity seem to be positively correlated (spatially) to FPR. There are some plausible explanations for this: ”

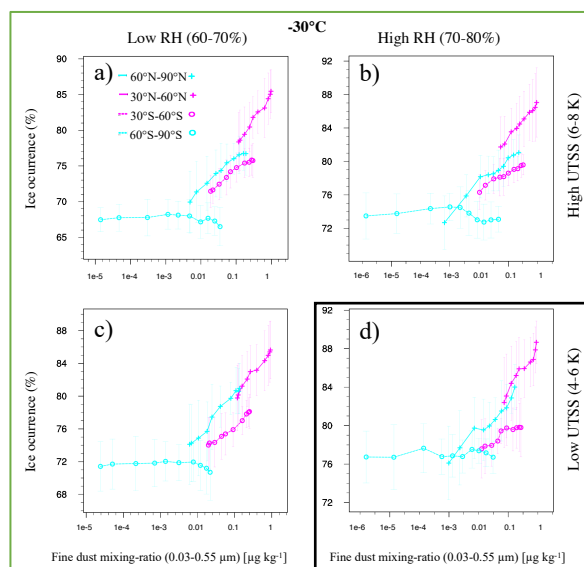
We have calculated the regression coefficient associated with the zonal averages and the  $30^{\circ}\times 1.875^{\circ}$  gridbox averages and we have also emphasized that we suggest only a large-scale correlation with the average updraft velocity. We have changed these lines as follows:

+(new Lines 375-380)“The time-averaged large-scale vertical velocity (MACC reanalysis) of the clouds studied is regionally correlated with the FPR at  $-15^{\circ}\text{C}$  — with a pearson correlation coefficient of 0.47 using zonal averages and of 0.31 using the  $30^{\circ}\times 1.875^{\circ}$  gridbox averages. Moreover, in another study, the spatial correlation between large-scale updraft velocity at 500 hPa was also found to be positively correlated (spatially) to the occurrence-frequency of ice clouds at  $-20^{\circ}\text{C}$  (Li et al., 2017a). In other words, both the dust mixing-ratio and the large-scale vertical velocity appear to be to some extent correlated (spatially) to the FPR. There are some plausible explanations for this: ”

- Given that a correlation with two different parameters of the FPR is studied, is it the case to carry out a partial correlation analysis?

The vertical velocity from the reanalysis is only a large-scale estimation and it may not coincide with the instant position of the clouds retrieved from the satellite products. For the same reason, LTSS and RH are better parameters to evaluate the possible influence of convection and constrain the influence of dynamics in the dust-cloud-phase correlation. Within this perspective, when we study the response of cloud-phase to different aerosol concentrations at constant LTSS and RH we fulfil the same objective that of a partial correlation analysis. As for the effect of dynamics (i.e., effect of updraft at a constant aerosol loading), this analysis has been already carried out in previous studies (e.g., Li et al., 2017) and is not the focus of this study.

- Lines 423 - 427: I agree with the statement provided by the authors though they should acknowledge that in the SH with low UTSS and high RH the positive correlation is much lower than in other conditions.



(Lines 423-428)—“ For dust mixing ratios between 0.1 and  $1.5 \mu\text{g kg}^{-1}$ , the cloud ice occurrence-frequency at  $-30^{\circ}\text{C}$  increase by about +5%. The highest increase is found for the northern latitudes. However, the results from the southern mid-latitudes contradict the notion that the INP activity of mineral dust is of secondary importance in the Southern Hemisphere due to low dust aerosol concentrations (Burrows et al., 2013; Kanitz et al., 2011). Nevertheless, recent studies have acknowledged that the importance of mineral dust in the southern latitudes still cannot be ruled out (Vergara-Temprado et al., 2017) ”

- Can the authors comment a bit more on this aspect?



We agree with the referee. However, because the difference is only evident for the southern mid-latitudes, it is very difficult to speculate about the reason behind it. Because this difference is not found in the NH, it is difficult to attribute the effect to the stability and humidity conditions. However, the correlation seems to vary little between the regimes, and therefore, it suggests that the positive correlation is consistent for the different cloud-forming conditions.

- Do the authors envisage a larger contribution in the SH of the homogenous nucleation than in other regions?

We consider that still much investigation is needed before taking a stand in this question. The relative contribution of homogeneous and heterogeneous freezing — and of the different INP types — is still a matter of debate (Barahona et al., 2017; Dietlicher et al., 2018), especially in the mixed-phase regime (temperature range 0°C to −42°C). Furthermore, even if mineral dust was not a dominant INP in the SH, other particles like marine organic aerosols could still represent important INP and influence cloud ice formation.

- Lines 435-439: these lines are too speculative, I'd honestly remove them.

(Lines 435-439)—“In general, for temperatures between −36°C and −9°C, higher fine-mode dust mixing-ratios are associated with an increasing cloud ice occurrence-frequency. The results suggest that only the lower static stability at −15°C has a strong influence on the relationship between mineral dust and cloud ice. This may be a consequence of the dynamic component of the atmospheric stability at lower temperatures (e.g., gravity waves), which is not included in the static stability parameter.”

Removed.

- Lines 441-443: Can these results be due to the purer nature of the dust in the SH compared to the NH, where it is often mixed to other aerosol types? In the discussion following to these lines, the authors mention the aged aerosol but never the effect of the aerosol mixing.

(Lines 441-443)—“However, against our expectations, for similar dust loadings the cloud ice occurrence-frequency at −15°C was higher in the Southern than in the Northern Hemisphere.”

- This is a very interesting point. Indeed, the higher concentrations of sulphate in the NH are believed to produce coating in dust aerosol and deactivate its freezing potential. It is therefore not difficult to imagine that mixing with other types of aerosols may cause a similar effect. We do mention in the text the potential role of biogenic aerosol mixed with dust aerosol, which would have an enhancing effect in the freezing potential. When we mentioned aged aerosol we refer mostly to the internal mixing (coating) of sulphate with dust aerosol. We have mentioned this explicitly.

+(new Line 509) “The ageing (e.g., internal mixing with sulfate or “coating”) of dust particles may also reduce the freezing efficiency of dust aerosol during the transport from low to high latitudes.”

- Line 502: among the significant number of factors contributing to the uncertainty affecting the presented analysis I'd add the limitation to consider only a specific type of cloud type, and only night time observations, as well as the effect of the electric charge of mineral dust particles.

- We have added these limitations to the list:

+(new Lines 542-549)

- Changes in dynamical forcing (e.g., updrafts) and **cloud regimes**
- Temperature changes after cloud glaciation (e.g., latent heat release)
- Ice sedimentation from above (cloud seeding), and INPs other than dust

- *Cloud vertical distribution within the studied temperature ranges*
  - *Turbulence favouring aerosol mixing and sub-grid temperature fluctuations*
  - *Differences in dust mineral composition, **electric charge** and/or size*
  - *Coatings (e.g. Sulfate) affecting aerosol solubility and freezing efficiency*
  - *Subsetting of the data (e.g., only night-time retrievals)”*
- Line 535 and following: in this section there are few sentences which are very speculative and though these are able put on the table the plethora of different interpretations to the presented data, at the same time, may be not always helpful to the users, also considering that this is not a research article. I suggest to shorten it or arrange in clearer way.
    - To improve the clarity, we have shortened these paragraphs as follows:
    -

+(new Lines 583-601)“ *In general, meteorological parameters have a larger impact on cloud properties than aerosols do (Gryspeerd et al., 2016). For example, different updraft regimes can change the aerosol-cloud interactions in warm clouds by an order of magnitude. Therefore, it is important to study how such meteorological parameters relate to the dust aerosol loading. With this purpose, Fig. 11 shows the mean relative humidity, cloud height and large-scale updraft at  $-15^{\circ}\text{C}$  for the different fine-mode dust mixing-ratio deciles and for the four latitude bands studied in Sect. 4.4. Firstly, the correlation between fine-mode dust mixing-ratio from the MACC reanalysis and the RH from the ERA-Interim reanalysis — weighted by cloud volume fraction — was found to be negative (Fig. 11a). We note that the RH from the ERA-Interim reanalysis represents the conditions at a large-scale and not the conditions at a specific location and the moment of the interaction between dust aerosol and supercooled cloud droplets. Still, this relationship is consistent with the intuition that dust is mostly associated with drier air masses. Second, The significant positive correlation found between dust aerosol mixing-ratio and the height of the isotherms (weighted by cloud volume fraction) points to an important source of uncertainty (Fig. 11b). This could be due to clouds being detected in a higher temperature bin after being glaciated at lower temperatures, thus erroneously suggesting an enhanced glaciation occurrence frequency at higher temperatures. Therefore, it is crucial for future studies to take into account this possibility when studying the occurrence of ice clouds at a certain isotherm. More details on the spatiotemporal variability of the cloud height can be found in the supplement (S12) to this article. Lastly, Fig. 11c shows a positive correlation between the fine-mode dust and the large-scale vertical velocity from the MACC reanalysis at  $-15^{\circ}\text{C}$ . Updrafts favour saturation over liquid water and therefore CCN activation, droplet growth and inhibition of the WBF (Wegener–Bergeron–Findeisen) process. Therefore, a positive dust-updraft correlation could lead to an underestimation of the dust-cloud-phase relationship.”*

Line 547: something missing in this sentence.

—“It is possible to find cases where the reanalysis and the detected have different temperatures.”

Meant was:

+“It is possible to find cases where the reanalysis and the detected **clouds** have different temperatures.”

However, we have removed this line after the previous suggestion to shorten this section.