# **Response to comments 1#**

Review of the manuscript revision by G. Zhang et al.

Overall, the authors have done well to address some of the referee comments, but not all. There are several major issues that remain, which prevent me from endorsing the manuscript for publication. For the comments that have not been addressed sufficiently, I have the original referee comment below in quotes, followed by my assessment of the changes and additional comments.

We would like to thank the referee again for his/her valuable comments to further improve the manuscript. We have addressed the specific comments in the sections below and made the appropriate revisions to the manuscript. The referee's comments are in the black text followed by our response in blue text.

Original comment: "For NH<sub>3</sub>, a positive correlation is observed between the number fraction of particles with NH4+ and NOC, while a negative correlation is observed between the relative peak areas of these compound classes. There is not a reasonable explanation given for this surprising and apparently contradictory behavior."

New comment: The authors' explanation for this observation still does not make sense. The authors state that ammonium is necessary for NOC formation, but that ammonium also inhibits NOC formation. Various factors contributing to this counterintuitive result are hypothesized (e.g., acidity, competition for gas-phase ammonia), but none are consistent with the results in Figure 3c and 3d. The discussion in the revised manuscript (revised manuscript with track changes lines 205 - 214) actually contradicts their finding, since ammonium does not inhibit NOC formation in any of the references cited in this added paragraph. Based on the other referee's comments, it seems there is a possibility that the results in Figures 3c and 3d stem from a measurement artifact (ionization efficiency changing with composition), rather than an actual physical/chemical process occurring in the atmosphere. The authors need to present a detailed, logical argument for this observation.

Thanks for pointing this out. Regarding the referee's concern, we also agree with

the comment that "ammonium also inhibits NOC formation" is not precise enough to pinpoint the observation results in Figure 3. It has been revised to "The formation of NOCs requires the presence of ammonium, but is most probably influenced by the relative amount of ammonium in individual particles" in the revised manuscript, which is believed to be more appropriate.

We understand that the reviewer still considered this observation as a counterintuitive result. As previously replied, we highlighted the contradictory correlation between the hourly detected number and RPAs of NOCs and ammonium as a surprising and interesting result. We think that the referee's confusion maybe mainly due to our unclear expressions on this issue. As replied to the referee 2#, this is unlikely due to measurement artifact. To make it clearer, we attempted to provide more discussion in the revised manuscript.

First, from the observational results in Figure 3d, we stated that the formation of NOCs requires the involvement of ammonium. It is easy to understand that ammonium is important in the formation of NOCs, from the high correlations between their hourly detected number. Consistently, NOCs and ammonium showed high internal mixing state in Figure 2.

Second, we provided additional explanations on the inverse correlation between RPAs of NOCs and ammonium as shown in Figure 3c. The negative correlation between their RPAs signifies that the formation of NOCs is most probably influenced by the relative amount of ammonium in individual particles. This is supported by our data, both from filter samples and individual particle analysis. There is a negative correlation between concentrations of WSON and  $NH_4^+$  for the filter samples (Fig. S6). It can be seen from Fig. S7 that lower RPAs of ammonium correspond to higher Nfs of ammonium that internally mixed with NOCs. Such an inverse correlation could also serve as evidence to explain the influence of the relative amount of ammonium on the formation of NOCs.

Furthermore, we discussed the theoretical possibility that the formation of NOCs is influenced by the relative amount of ammonium. We showed that this is theoretically possible since the formation of NOCs may be enhanced by particle acidity (Miyazaki

et al., 2014; Aiona et al., 2017; Nguyen et al., 2012), which is substantially affected by the amount of ammonium. This could also be supported by our data, as shown in Figure S6, that the internally mixed ammonium and NOCs particles had higher relative acidity than the ammonium-containing particles without NOCs. It is also noted that particle acidity could play a significant role in the gas-to-particle partitioning of aldehydes (Herrmann et al., 2015; Liggio et al., 2005; Gen et al., 2018; De Haan et al., 2018; Kroll et al., 2005), which are precursors for the formation of the oxidized organics.

Original comment: "NOx is completely ruled out as a contributor to NOC formation on the basis of poor (or no) correlations between NOC and NOx. However, this is a misinterpretation of the data. Many factors (different removal processes and lifetimes of particles vs. gasses, primary vs. secondary species, etc.) could contribute to a lack of correlation even if NOx did contribute to NOC formation."

New comment: First of all, the authors used the referee's comment word-for-word in their revised manuscript (lines 460-462 in the track changes version). This is inappropriate, and constitutes plagiarism of this sentence. Clearly, this should be changed.

Thanks for pointing this out. We have changed the sentence to "A lack of correlation could probably be attributed to that NOx affects the formation of NOCs in various ways (e.g., peroxy radical chemistry in VOCs oxidation mechanisms and formation of nitrate radicals) (Xu et al., 2015; Zhang et al., 2018), and thus may not linearly contribute to the formation of NOCs."

The above issue aside, I still do not believe this comment was adequately addressed in the revision. There is too much emphasis on simple linear correlations, when that is not expected for the chemistry in this system. For example, NOx controls the branching of VOC reactions, which will in turn affect product distributions, including NOCs; however, this will not (in most cases) result in a linear correlation between NOx and NOCs. Likewise, NOx affects nitrate radical formation, which can form NOCs, but a linear relationship between NOx and NOCs will not necessarily occur even if this is the predominant pathway for NOC production. Further, the references cited in lines 451-453 (track changes version) to support their position do not show linear correlations between NOx and NOC formation in systems representative of a polluted urban atmosphere. Therefore, the possible role of NOx in NOC formation is not accurately described in the manuscript revision.

Thanks for the referee's constructive comment. It greatly helps us to refine our discussion on the possible contribution of NOx pathways to the formation of NOCs. We agree with the comment that simple linear correlations may not be expected for the chemistry involving NOx and NOCs. This is because that NOx affects the formation of NOCs in various ways (e.g., peroxy radical chemistry in VOCs oxidation mechanisms and formation of nitrate radicals) (Xu et al., 2015; Zhang et al., 2018). Besides, the related products or intermediates were not available to elucidate such complex chemical processes in the present study. Therefore in the revised manuscript, we mentioned that NOx may play a certain role in the conversion of oxidized organics to NOCs, and yet it cannot be quantified in the present study. This is supported by the relationship between peak ratios of NOCs/oxidized organics and NOx (Fig. 6), showing that the ratios generally increase with an increasing level of NOx ( $R^2 = 0.18$ , p < 0.01). It is noted that low correlation coefficients ( $R^2 = 0.02 - 0.13$ ) between NOx and NOCs might not indicate a limited contribution of NOx to the formation of NOCs. A lack of correlation could probably be interpreted by the fact that NOx affects the formation of NOCs in various ways (e.g., peroxy radical chemistry in VOCs oxidation mechanisms and formation of nitrate radicals) (Xu et al., 2015; Zhang et al., 2018), and thus may not linearly contribute to the formation of NOCs.

Finally, the revised manuscript needs editing for grammar and language.

Thanks for the comment. We have carefully examined and corrected possible technical errors.

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1	High secondary formation of nitrogen-containing organics (NOCs) and its
2	possible link to oxidized organics and ammonium
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20	Hi	ghlights
21	•	Nitrogen-containing organics (NOCs) were highly internally mixed with photochemically
22		produced secondary oxidized organics
23	•	NOCs could be well predicted by the variations of these oxidized organics and ammonium
24	•	Higher relative humidity and NOx may facilitate the conversion of these oxidized organics
25		to NOCs

### 26 Abstract

27 Nitrogen-containing organic compounds (NOCs) substantially contribute to light-light-28 absorbing organic aerosols, although the atmospheric processes responsible for the secondary 29 formation of these compounds are poorly understood. In this study, seasonal atmospheric 30 processing of NOCs was were investigated by single single-particle mass spectrometry in urban 31 Guangzhou from 2013-2014. The relative abundance of NOCs is found to be strongly enhanced 32 when internally mixed with the photochemically produced secondary oxidized organics (i.e., 33 formate, acetate, pyruvate, methylglyoxal, glyoxylate, oxalate, malonate, and succinate) and 34 ammonium. In additionBesides, both the hourly detected particle number and relative 35 abundance of NOCs are highly correlated with those of secondary oxidized organics and 36 ammonium. It is therefore hypothesized that secondary formation of NOCs most likely links to 37 the oxidized organics and ammonium. Results from both multiple linear regression analysis and 38 positive matrix factorization analysis further show that the relative abundance of NOCs could 39 be well predicted ( $R^2 > 0.7$ , p < 0.01) by the oxidized organics and ammonium. 40 Interestingly, the relative abundance of NOCs is inversely correlated with ammonium, 41 whereas their number fractions are positively correlated. This result suggests that although the 42 formation of NOCs does require the involvement of NH3/NH4 , the relative amount of 43 ammonium may have a negative effect. Higher humidity and NOx likely facilitate tThe 44 conversion of oxidized organics to NOCs is likely facilitated by higher humidity and NOx. Due 45 to the relatively high oxidized organics and NH<sub>3</sub>/NH<sub>4</sub><sup>4</sup>, the relative contributions of NOCs in

46 summer and autumn were higher than those in spring and winter. To the best of our knowledge,

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47	this is the first direct field observation study reporting a close association between NOCs and
48	both oxidized organics and ammonium. These findings have substantial implications for the
49	role of ammonium in the atmosphere, particularly in models that predict the evolution and
50	deposition of NOCs.

- 51
- 52 Keywords: nitrogen-containing organic compounds, individual particles, oxidized organics,

53 ammonium, mixing state, single-single-particle mass spectrometry

## 55 1 Introduction

56 Organic aerosols that strongly absorb solar radiation are referred to as brown carbon 57 (BrC), capable of a comparable level of light absorption in the spectral range of near-58 ultraviolet (UV) light as black carbon (Andreae and Gelencser, 2006; Feng et al., 2013; Yan 59 et al., 2018). Nitrogen-containing organic compounds (NOCs) substantially contribute to the 60 pool of BrC (Feng et al., 2013; Mohr et al., 2013; Li et al., 2019), and have a major 61 significant effect on atmospheric chemistry, human health and climate forcing (Noziere et 62 al., 2015; Kanakidou et al., 2005; Shrivastava et al., 2017; De Gouw and Jimenez, 2009). 63 The particulate organic nitrogen accounts for a large fraction of total airborne nitrogen 64 (~30%), although the proportion exhibits a high variability temporally and spatially, and 65 therefore has an influence on both regional and global N deposition (Neff et al., 2002; Shi et 66 al., 2010; Cape et al., 2011). However, the sources, evolution, and optical properties of NOCs 67 remain unclear and contribute significantly to uncertainties in the estimation of their impacts 68 on the environment and climate (Laskin et al., 2015; Feng et al., 2013). 69 NOCs are ubiquitous components of atmospheric aerosols, cloud water and rainwater 70 (Altieri et al., 2009; Desyaterik et al., 2013; Laskin et al., 2015), spanning a wide range of 71 molecular weights, structures and light absorption properties (Lin et al., 2016). Emissions of

72 primary NOCs have been attributed to biomass burning, coal combustion, vehicle emissions,

biogenic production and soil dust (Laskin et al., 2009; Desyaterik et al., 2013; Sun et al.,
2017; Mace et al., 2003; Rastogi et al., 2011; Wang et al., 2017). A growing body of evidence

75	from laboratory studies suggests that secondary NOCs may be produced in gas phase,
76	aerosol, and clouds. Maillard reactions involving mixtures of atmospheric aldehydes (e.g.,
77	methylglyoxal/glyoxal) and ammonium/amines are of particular interests (e.g., Hawkins et
78	al., 2016; De Haan et al., 2017; De Haan et al., 2011). A significant portion of NOCs may
79	also be derived from the heterogeneous ageing of secondary organic aerosol (SOA) with
80	$\rm NH_3$ / $\rm NH_4^{\scriptscriptstyle +}$ (Liu et al., 2015; Laskin et al., 2015). Mang et al. (2008) proposed that even
81	trace levels of ammonia may be sufficient to form NOCs via this pathway. In addition, gas
82	gas-phase formation of NOCs through interaction between volatile organic hydrocarbons
83	and NO <sub>x</sub> and other oxidations, followed by condensation, may have <u>a</u> potential contribution
84	(Fry et al., 2014; Stefenelli et al., 2019; Lehtipalo et al., 2018).
85	The secondary formation of NOCs is especially prevalent in environments experiencing
86	high anthropogenic emissions (Yu et al., 2017; Ho et al., 2015), although further studies are
87	required to comprehensively establish the formation mechanisms comprehensively. A major
88	obstacle is that organic and inorganic matrix effects have a profound impact on the chemistry
89	of organic compounds in bulk aqueous particles and particles undergoing drying (El-Sayed

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et al., 2015; Lee et al., 2013). While real-time characterization studies remain a challenge

due to the extremely complex chemical nature of NOCs, establishing this data along with

the co-variation of NOCs with other chemical components would help to identify the sources

and evolution of NOCs. Using single-particle aerosol time-of-flight mass spectrometry,

Wang et al. (2010) observed that the widespread occurrence of NOCs was closely correlated

with particle acidity in the atmosphere of Shanghai (China). In addition, real-time

96	measurements of the atmosphere in New York (US) by aerosol mass spectrometry indicated	
97	a positive definite link between the age of organic species and the N/C ratio (Sun et al., 2011).	<b>带格式的:</b> 字体: Times New Roman
98	Further in-depth studies are required to identify the role of formation conditions (e.g.,	
99	relative humidity (RH) and pH) for secondary NOCs (Aiona et al., 2017; Nguyen et al.,	
100	2012). In the present study, the mixing state of individual particles were was investigated,	
101	involving NOCs, oxidized organics, and ammonium, based on on-line seasonal observations	
102	using a single particle aerosol mass spectrometry (SPAMS). Our findings show that the	
103	formation of NOCs is significantly linked to oxidized organics and $\mathrm{NH}_4^{\scriptscriptstyle +}$ , which has	
104	important environmental implications for assessing the impact and fate of these compounds.	
105		
106	2 Methods	
100		
107	2.1 Field measurements	
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107 	2.1 Field measurements	<b>带格式的:</b> 字体: Times New Roman
107 108	2.1 Field measurements Sampling was performed constructed at the Guangzhou Institute of Geochemistry, a	<b>带格式的:</b> 字体: Times New Roman
107 108 109	<b>2.1 Field measurements</b> Sampling was <u>performed-constructed</u> at the Guangzhou Institute of Geochemistry, a representative urban site in Guangzhou (China), a megacity in the Pearl River Delta (PRD)	<b>带格式的:</b> 字体: Times New Roman
107 108 109 110	2.1 Field measurements Sampling was <u>performed-constructed</u> at the Guangzhou Institute of Geochemistry, a representative urban site in Guangzhou (China), a megacity in the Pearl River Delta (PRD) region. <u>The size and chemical composition of individual particles were obtained by the</u>	<b>带格式的:</b> 字体: Times New Roman <b>带格式的:</b> 字体: Times New Roman
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2014), summer (13/06 to 16/07 2013), autumn (26/09 to 19/10 2013), and winter (15/12 to
25/12 2013). The total measured particle numbers and mean values for meteorological data
and gaseous pollutants, are outlined for each season in Table S1 and were described in a
previous publication (Zhang et al., 2019).

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#### 122 2.2 SPAMS data analysis

123 Fragments of NOCs were identified according to the detection of ion peaks at m/z -26 124 [CN]<sup>-</sup> or -42 [CNO]<sup>-</sup>, generally due to the presence of C-N bonds (Silva and Prather, 2000; 125 Zawadowicz et al., 2017; Pagels et al., 2013). Laboratory produced C-N bonds compounds 126 from bulk solution-phase reactions between the representative oxidized organics (i.e., 127 methylglyoxal) and ammonium sulfate was used to confirm the generation of ion peaks at 128 m/z -26 [CN]<sup>-</sup> and/or -42 [CNO]<sup>-</sup> using SPAMS (Fig. S1). Thus, the NOCs herein may refer 129 to complex nitrated organics such as organic nitrates, nitro-aromatics, nitrogen heterocycles. 130 and polyphenols. Unfortunately, how well [CN]<sup>-</sup>/ [CNO]<sup>-</sup> ions could represent NOCs cannot 131 be quantified, although they were the most commonly reported NOCs peaks by single-single-132 particle mass spectrometry (Silva and Prather, 2000; Zawadowicz et al., 2017; Pagels et al., 133 2013). In the present study, [CN]<sup>-</sup>/ [CNO]<sup>-</sup> ions are among the major peaks detected by the 134 SPAMS (Fig. 1). A rough estimate from the peak area ratio of [CN]<sup>-</sup>/[CNO]<sup>-</sup> ions and the 135 most likely NOCs fragments (i.e., various amines, and an entire series of nitrogen-containing 136 cluster ions  $C_nN^-$ , n = 1, 2, 3, ...) (Silva and Prather, 2000) shows that  $[CN]^- / [CNO]^-$  ions 137 may represent more than 90% of these NOCs peaks. The number fractions (Nfs) of particles

that contained NOCs ranged from 56-59% across all four seasons (Table S1). The number of detected NOCs-containing particles distributing along their vacuum aerodynamic diameter ( $d_{va}$ ) is shown in Fig. S2. Most of the detected NOC-containing particles had a  $d_{va}$ in a range of 300-1200 nm.

142 A representative mass spectrum for NOCs-containing particles is shown in Fig. 1. 143 Dominant peaks in the mass spectrum were 39 [K]<sup>+</sup>, 23 [Na]<sup>+</sup>, nitrate (-62 [NO<sub>3</sub>]<sup>-</sup> or -46 144 [NO<sub>2</sub>]<sup>-</sup>), sulfate (-97 [HSO<sub>4</sub>]<sup>-</sup>), organics (27 [C<sub>2</sub>H<sub>3</sub>]<sup>+</sup>, 63 [C<sub>5</sub>H<sub>3</sub>]<sup>+</sup>, -42 [CNO]<sup>-</sup>, -26 [CN]<sup>-</sup>), 145 ammonium (18  $[NH_4]^+$ ) and carbon ion clusters ( $C_n^{+/-}$ , n = 1, 2, 3,...). NOCs-containing 146 particles were internally mixed with various oxidized organics, represented as formate at m/z 147 -45 [HCO<sub>2</sub>]<sup>-</sup>, acetate at m/z -59 [CH<sub>3</sub>CO<sub>2</sub>]<sup>-</sup>, methylglyoxal at m/z -71 [C<sub>3</sub>H<sub>3</sub>O<sub>2</sub>]<sup>-</sup>, glyoxylate 148 at m/z -73  $[C_2HO_3]^-$ , pyruvate at m/z -87  $[C_3H_3O_3]^-$ , malonate at m/z -103  $[C_3H_3O_4]^-$  and 149 succinate at  $m/z - 117 [C_4H_5O_4]^-$  (Zhang et al., 2017; Zauscher et al., 2013; Lee et al., 2003). 150 These oxidized organics showed their pronounced diurnal trends with afternoon maximum\_3 151 and were highly correlated (r = 0.72 - 0.94, p < 0.01) with each other. Therefore, they were 152 primarily attributed to secondary oxidized organics from photochemical oxidation products 153 of various volatile organic compounds (VOCs) (Paulot et al., 2011; Zhao et al., 2012; Ho et 154 al., 2011), and the details can be found in our previous publication (Zhang et al., 2019). More 155 information on the seasonal variation range of the Nfs of oxidized organics, ammonium and 156 NOCs is presented in Fig. S3. 157 Hourly mean Nfs and relative peak areas were applied herein to indicate the variations

158 of aerosol compositions in individual particles. Even though advances have been made in

159	the quantification of specific chemical species for individual particles based on their
160	respective peak area information, it is still quite a challenge for SPAMS to provide
161	quantitative information on aerosol components mainly due to matrix effects, incomplete
162	ionization and so forth (Qin et al., 2006; Jeong et al., 2011; Healy et al., 2013; Zhou et al.,
163	2016). Despite of this, the variation of relative peak area should be a good indicator for the
164	investigation of atmospheric processing of various species in individual particles (Wang et
165	al., 2010; Zauscher et al., 2013; Sullivan and Prather, 2007; Zhang et al., 2014).
166	
167	3 Results and Discussion
168	3.1 Evidence for the formation of NOCs from oxidized organics and ammonium
169	Figure 2 shows the seasonal variations in Nfs of the oxidized organics and ammonium,
170	which were internally mixed with NOCs. On average, more than 90% of the oxidized
171	organics and 65% of ammonium (except spring) were found to be internally mixed with
172	NOCs (Fig. S4). Regarding that the Nfs of NOCs relative to all the measured particles was
173	~60%, it could be concluded that NOCs were enhanced with the presence of oxidized
174	organics and ammonium, with the enhancement associated with oxidized organics being the
175	most pronounced.
176	A strong correlation between both the Nfs and relative peak areas (RPAs) of NOCs and
177	oxidized organics further demonstrates their close associations, as shown in Fig. 3.
178	Compared with the oxidized organics, the Nfs of ammonium-containing particles internally

enhancement mixing of NOCs with ammonium. A positive correlation ( $R^2 = 0.50$ , p < 0.01) is observed between the hourly detected number of NOCs and ammonium. It is worth noting that a negative correlation ( $R^2 = 0.55$ , p < 0.01) is obtained between the hourly average RPAs of NOCs and ammonium (Fig. 3).

184 Based on both the enhancement of NOCs and the high correlations with oxidized 185 organics and ammonium, it is hypothesized that interactions between oxidized organics and 186 ammonium contributed to the observed NOCs. Actually, The formation of NOCs from 187 ammonium and carbonyls has been confirmed in several laboratory studies (Sareen et al., 188 2010; Shapiro et al., 2009; Noziere et al., 2009; Kampf et al., 2016; Galloway et al., 2009). 189 Secondary organic aerosols (SOA) produced from a large group of biogenic and 190 anthropogenic VOCs can be further aged by NH<sub>3</sub>/NH<sub>4</sub><sup>+</sup> to generate NOCs (Nguyen et al., 191 2012; Bones et al., 2010; Updyke et al., 2012; Liu et al., 2015; Huang et al., 2017). In a 192 chamber study, the formation of NOCs is enhanced in an NH3-rich environment (Chu et al., 193 2016). While such chemical mechanisms might be complexcomplicated, the initial steps 194 generally involve reactions forming imines and amines, which can further react with 195 carbonyl SOA compounds to form more complex products (e.g., oligomers/BrC) (Laskin et 196 al., 2015).

To verify this hypothesis, multiple linear regression analysis is performed to test how well the RPAs of NOCs could be predicted by the oxidized organics and ammonium. As expected, there is a close association ( $R^2 = 0.71$ , p < 0.01) between the predicted RPAs and the observed values of NOCs (Fig. 4), which supports this hypothesis. An obvious noticeable

201	improvement in R <sup>2</sup> implies that a model that uses both oxidized organics and ammonium to
202	predict RPAs of NOCs is substantially better than one that uses only one predictor (either
203	oxidized organics or ammonium in Fig. 3). The result indicates that interactions involving
204	oxidized organics and ammonium could explain over half of the observed variations in
205	NOCs in the atmosphere of Guangzhou. A fraction of the unaccounted NOCs could be due
206	to primary emissions and other formation pathways. This hypothesis could also be supported
207	by the <u>a</u> similar pattern of diurnal variation observed for NOCs and oxidized organics (Fig.
208	S5), although there is a slight lag for the NOCs. Such <u>a</u> diurnal pattern is similar to those
209	observed in Beijing and Uintah (Yuan et al., 2016; Zhang et al., 2015). Notably, such a
210	diurnal pattern of secondary NOCs is <u>effectively-adequately</u> modelled when the production
211	of NOCs via carbonyls and ammonium is included (Woo et al., 2013). In addition to possible
212	photo-bleaching (Zhao et al., 2015), the lower contribution of NOCs during the daytime may
213	be partly explained by the lower RH, as discussed in section 3.2.
214	Interestingly, the relationship between NOCs and ammonium is distinctly different from
215	the relationship between NOCs and oxidized organics (Fig. 3). This implies that the
216	controlling factors on the formation of NOCs from ammonium are different from oxidized
217	organics. On the one hand, the positive correlation between the detected numbers reflects
218	that the formation of NOCs does require the participant of $\rm NH_3/\rm NH_4^{\scriptscriptstyle +}$ , consistent with the
219	enhancement of NOCs in ammonium-containing particles (Fig. 2) discussed above. On the
220	other hand, the negative correlation between the RPAs signifies that particles with higher
221	relative ammonium content may inhibit the formation of NOCs the formation of NOCs is

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222	most probably influenced by the relative amount of ammonium in individual particles. Such
223	influence could also be supported by our data, both from filter samples and individual
224	particle analysis. There is a negative correlation between concentrations of WSON and NH
225	$\frac{1}{4}$ for the filter samples (Fig. S6). It can be seen from Fig. S7 that lower RPAs of ammonium
226	correspond to higher Nfs of ammonium that internally mixed with NOCs. Such an inverse
227	correlation could also serve as evidence to explain the influence of the relative amount of
228	ammonium on the formation of NOCs. Consistently, there is a negative correlation between
229	concentrations of WSON and NH <sup>+</sup> in filter samples (Fig. S6).
230	The influence of relative ammonium amount on the formation of NOCs This is
231	supported by the inverse correlation between that Nfs of ammonium that internally mixed
232	with NOCs and the RPAs of ammonium (Fig. S7). This is also theoretically possible since
233	the formation of NOCs may be influenced-enhanced by particle acidity (Miyazaki et al.,
234	2014; Aiona et al., 2017; Nguyen et al., 2012), which is substantially affected by the
235	abundance of ammonium. Consistently, higher relative acidity was observed for the
236	internally mixed ammonium and NOCs particles, compared to ammonium-containing
237	particles without NOCs (Fig. S6) and thus may influence the formation of NOCs (Fig. S7).
238	Particle acidity could also play a significant role in the gas-to-particle partitioning of
239	aldehydes (Herrmann et al., 2015; Liggio et al., 2005; Gen et al., 2018; De Haan et al., 2018;
240	Kroll et al., 2005), precursors for the formation of oxidized organics. Consistently, higher
241	relative acidity was observed for the internally mixed ammonium and NOCs particles,
242	compared to ammonium-containing particles without NOCs (Fig. S6), and thus may

243	influence the formation of NOCs (Fig. S7). However, the higher relative acidity might also
244	be a result of NOCs formation. A model simulation shows that after including the chemistry
245	of SOA ageing with NH <sub>3</sub> , an increase in aerosol acidity would be expected due to the
246	reduction in ammonium (Zhu et al., 2018). It is also noted that the particle acidity is roughly
247	estimated by the relative abundance of ammonium, nitrate, and sulfate in individual particles
248	(Denkenberger et al., 2007), and thus may not be representative of actual aerosol acidity or
249	pH (Guo et al., 2015; Hennigan et al., 2015; Murphy et al., 2017). In addition, ammonia in
250	the gas phase is also efficient at producing NOCs (Nguyen et al., 2012), which may play-a
251	an intricate complex role in the distribution of ammonium and NOCs in the particulate phase.
252	The formation of ammonium and NOCs would compete for ammonia, which may also
253	potentially result in the negative correlation between the RPAs of NOCs and ammonium.
254	Unfortunately, such a role remains unclear since the variations of ammonia were not
255	available in the present study.
256	
257	3.2 Factors contributing to the NOCs resolved by positive matrix factorization (PMF)
258	analysis
259	Figure 5 presents the PMF factor profiles obtained from the PMF model analysis
260	(detailed information is provided in the SI) (Norris et al., 2009) and their diurnal variations.
261	Around 75% of NOCs could be well explained by two factors, with 33% of the PMF resolved
262	NOCs mainly associated with ammonium and carbonaceous ion peaks (ammonium factor),

263 while 59% were mainly associated with oxidized organics (oxidized organics factor). The

264	explained fraction of NOCs by the ammonium and oxidized organic factors is consistent
265	with the linear regression analysis. In addition Furthermore, PMF analysis provided
266	information on the factor contribution and diurnal variations, which may help explain the
267	seasonal variations and processes of NOCs. The ammonium factor showed a diurnal
268	variation pattern peaking during the early morning, which is consistent with the diurnal
269	variation in RH (Zhang et al., 2019). This factor contributed to ~80% (Fig. S8) of the PMF
270	resolved NOCs during spring with the highest RH (Table S1), whereas the oxidized organics
271	factor dominated (> 80%) in summer and fall. In winter, these two factors similarly
272	contributed (~40%). This may Variation of the ammonium factor indicate may reflect a
273	potential role of aqueous pathways in the formation of NOCs, particularly during spring.
274	Differently, the oxidized organics factor showed a pattern of diurnal variation, increasing
275	from morning hours and peaking overnight, which may correspond to the photochemical
276	production of oxidized organics and followed interactions with condensed ammonium. This
277	pathway may explain the slightly late peaking of NOCs compared to oxidized organics, as
278	ammonium condensation is favorable overnight (Hu et al., 2008). While there were
279	similarities in the fractions of oxidized organics in the oxalate factor and the oxidized
280	organics factor, they only contributed to 8% of the PMF resolved NOCs in the oxalate factor,
281	which contained ~80% of the PMF resolved oxalate. As previously discussed, these oxidized
282	organics are also precursors for the formation of oxalate (Zhang et al., 2019). Therefore, the
283	PMF results suggest that there are two competitive pathways for the evolution of these
284	oxidized organics. Some oxidized organics formed from photochemical activities were

285	further oxidized to oxalate, resulting in a diurnal pattern of variation with concentration
286	peaks during the afternoon (Fig. 5), while others interact with $\rm NH_3/\rm NH_4^{\scriptscriptstyle +}$ to form NOCs,
287	peaking during the nighttime. However, the controlling factors for these pathways could not
288	be determined in the present study. The unexplained NOCs (~25%) might be linked to the
289	primary emissions, such as biomass burning (Desyaterik et al., 2013). It could be partly
290	supported by the presence of potassium and various carbon ion clusters ( $C_n^{*\!/}$ , n = 1, 2, 3,)
291	in the mass spectrum of NOCs-containing particles (Fig. 1).

## 293 3.3 Seasonal variations in the observed NOCs

294 There is a clear an evident seasonal variation of NOCs, with higher relative contributions 295 during summer and autumn (Figs. 3 and 4), mainly due to the variations in oxidized organics 296 and NH<sub>3</sub>/NH<sup>+</sup><sub>4</sub>. In this region, a larger-more considerable contribution from secondary 297 oxidized organics is typically observed during summer and autumn (Zhou et al., 2014; Yuan 298 et al., 2018). The seasonal maximum NH3 concentrations have also been reported during the 299 warmer seasons, corresponding to the peak emissions from agricultural activities and high 300 temperatures, while the low NH3 concentrations observed in colder seasons may be 301 attributed to gas-to-particle conversion (Pan et al., 2018; Zheng et al., 2012). Such seasonal 302 variation in NOCs is also obtained in a model simulation, showing that the conversion of 303 NH3 into NOCs would result in a significantly higher reduction of gas-phase NH3 during 304 summer (67%) than winter (31%), due to the higher NH<sub>3</sub> and SOA concentrations present 305 in the summer (Zhu et al., 2018). More primary NOCs may also be present during summer

306	and autumn in the present study, due to the additional biomass burning activities in these
307	seasons (Chen et al., 2018; Zhang et al., 2013).
308	The seasonal variations of NOCs can be adequately explained by the variations in
309	concentrations of oxidized organics and ammonium (Fig. 4), although the hourly variations
310	during each season are not well explained, as indicated by the lower $R^2$ values (Table S2).
311	The correlation coefficients $(R^2)$ range from 0.24 to 0.57 for inter-seasonal variations.
312	During spring, NOCs exhibits a limited dependence on oxidized organics (Figs. 3a and 3b),
313	while during summer, the hourly detected number of NOCs shows a limited dependence on
314	ammonium (Fig. 3d). These seasonal dependences of NOCs are consistent with results can
315	be explained by the PMF results, showing that the ammonium factor explained ~80% of the
316	predicted NOCs during spring, while the oxidized organics factor dominantly contributed to
317	the predicted NOCs during warmer seasons (Fig. S8). A detailed discussion of this issue is
318	provided in the SI.
319	
320	3.4 Influence of RH and NOx
321	The influence of RH on RPAs of NOCs and peak ratios of NOCs/oxidized organics, are
322	shown in Fig. 6. While NOCs do not show a clear dependence on RH, the ratio of NOCs to
323	the oxidized organics shows an apparenta clear increase towards higher RH. This finding is
324	consistent with the observations reported by Xu et al. (2017), in which the N/C ratio
325	significantly increases as a function of RH in the atmosphere of Beijing. In additionBesides,
326	the diurnal variations of NOCs with peaks values around 20:00 are also similar to those

327	reported by Xu et al. (2017). The peak ratios of NOCs/oxidized organics are more obviously
328	enhanced when RH is higher than 40%. These findings imply that aqueous-phase processing
329	likely plays an a substantial important role in the formation of NOCs. Significant changes in
330	RH, such as during the evaporation of water droplets, have been reported to facilitate the
331	formation of NOCs via $NH_3/NH_4^+$ and SOA (Nguyen et al., 2012). In addition, an increase
332	in RH would improve the uptake of $NH_3$ and <u>the</u> formation of $NH_4^*$ , which also contributes
333	to the enhancement of NOCs. However, the relatively weak correlation ( $R^2 = 0.27, p < 0.01$ )
334	between the peak ratios and RH, reflect the complex influence of RH on the formation of
335	NOCs (Xu et al., 2017; Woo et al., 2013).
336	One may expect that NOCs were are formed through the interactions between NOx and
337	oxidized organics in the gas phase, followed by condensation (Fry et al., 2014; Lehtipalo et
338	al., 2018). Similar to that observed for RH, while NOCs do not show a clear dependence on
339	NOx (Fig. 6c, $R^2 = 0.02-0.13$ ), the ratio of NOCs to the oxidized organics shows a clear
340	increasing trend towards higher NOx (Fig. 6d, $R^2 = 0.18$ , $p < 0.01$ ). Low correlation
341	coefficients ( $R^2 = 0.02 - 0.13$ ) between NOCs and NOx likely indicates limited contribution
342	of this pathways to the observed NOCs. We have also included an analysis on the
343	relationship between peak ratios of NOCs/oxidized organics and NOxThis indicates that
344	NOx may play a certain role in the conversion of oxidized organics to NOCs, and yet it
345	cannot be quantified in the present study. It is also noted that low correlation coefficients
346	between NOx and NOCs might not indicate a limited contribution of NOx to the formation
347	of NOCs. NOx affects the formation of NOCs in various ways (e.g., peroxy radical chemistry

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348	in VOCs oxidation mechanisms and formation of nitrate radicals) {Xu, 2015
349	#20234}{Zhang, 2018 #22855}, and thus may not linearly contribute to the formation of
350	NOCs. Peak area ratios of NOCs/oxidized organics generally increases with increasing level
351	of NOx (Fig. 6), but still with relatively weak correlation ( $\mathbb{R}^2 = 0.18, p < 0.01$ ). An inclusion
352	of both NOx and RH in the above linear regression model (NOCs versus the oxidized
353	organics and ammonium) does not improve the prediction of NOCs ( $R^2 = 0.71, p < 0.01$ ).
354	However, it is also noted that many factors (e.g., different removal processes and lifetimes
355	of particles vs. gasses, primary vs. secondary species, etc.) could contribute to a lack of
356	strong correlation even if NOx did contribute to NOC formation. (Xu et al., 2015)
357	

358 **3.5** Atmospheric implications and limitation

359 In this study, we showed that in an urban megacity area, secondary NOCs were 360 significantly contributed by the heterogeneous ageing of oxidized organics with  $NH_3/NH_4^*$ , 361 providing valuable insight into SOA aging mechanisms. In particular, the effects of NH<sub>3</sub>/NH 362 <sup>+</sup><sub>4</sub> on SOA or BrC formation remain relatively poorly understood. In the PRD region, it has 363 been shown that oxygenated organic aerosols (OOA) account for more than 40% of the total 364 organic mass (He et al., 2011), with high concentrations of available gaseous carbonyls (Li 365 et al., 2014). Therefore, it is expected that over half of all water\_-soluble NOCs in this region 366 might link to secondary processing (Yu et al., 2017). Furthermore, secondary sources have 367 been found to contribute significantly to NOCs related BrC in Nanjing, China (Chen et al., 368 2018). The results presented herein also suggest that the production of NOCs might be

369	effectivelyadequately estimated by their correlation with secondary oxidized organics and
370	ammonium. The effectiveness of correlation-based estimations needs to be examined in
371	other regions before being generally applied in other environments. However, this approach
372	may provide valuable insights into investigations into of NOCs using atmospheric
373	observations. In contrast, it has previously been reported that a positive correlation exists
374	between WSON and ammonium (Li et al., 2012), indicating similar anthropogenic sources.
375	This divergence could be mainly attributed to varying contributions of primary sources and
376	secondary processes to the observed NOCs. Possible future reductions in anthropogenic
377	emissions of ammonia may reduce particle NOCs. Understanding the complex interplay
378	between inorganic and organic nitrogen is an important essential part of assessing the global
379	nitrogen cycling.

380 Moise et al. (2015) proposed that with high concentrations of reduced nitrogen 381 compounds, high photochemical activity, and frequent changes in humidity, BrC formed via 382 NH<sub>3</sub>/NH<sub>4</sub><sup>+</sup> and SOA may become a dominant contributor to aerosol absorption, specifically 383 in agricultural and forested areas. However, this study suggests that even in typical urban 384 areas, BrC formation via  $NH_3/NH_4^+$  and SOA should not be neglected. In particular, SOA 385 was found to account for 44 - 71% of the organic mass in megacities across China (Huang 386 et al., 2014), with NH3 concentrations in urban areas comparable with those from 387 agricultural sites and 2- or 3-fold those of forested areas in China (Pan et al., 2018). 388 Additionally, the acidic nature of particles in these regions would also be also favorable for 389 the formation of NOCs (Guo et al., 2017; Jia et al., 2018). Considering the formation of

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390 NOCs from the uptake of  $NH_3$  onto SOA particles, Zhu et al. (2018) suggested that this 391 mechanism could have a significant impact on the atmospheric concentrations of  $NH_3/NH_4^+$ 392 and  $NO_3$ .

393

#### 394 5 Conclusions

395 This study investigated the processes contributing to the seasonal formation of NOCs, 396 involving ammonium and oxidized organics in urban Guangzhou, using single-single-397 particle mass spectrometry. This is the first study to provide direct field observation results 398 to confirm that the variation of NOCs correlate well and are strongly enhanced internal 399 mixing with secondary oxidized organics. These findings highlight the possible formation 400 pathway of NOCs through the ageing of secondary oxidized organics by NH<sub>3</sub>/NH<sub>4</sub><sup>+</sup> in 401 ambient urban environments. A clear pattern of seasonal variation in NOCs was observed, 402 with higher relative contributions in summer and autumn as compared to spring and winter. 403 This seasonal variation was well predicted by multiple linear regression model analysis, 404 using the relative abundance of oxidized organics and ammonium as model inputs. More 405 than 50% of NOCs could be explained by the interaction between oxidized organics and 406 ammonium. The production of NOCs through such processes were was facilitated by 407 increased humidity and NOx. These results extend our understanding of the mixing state and 408 atmospheric processing of particulate NOCs, as well as having important-substantial 409 implications for the accuracy of models predicting the formation, fate, and impacts of NOCs 410 in the atmosphere.

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412	Author contribution	
413	GHZ and XHB designed the research (with input from WS, LL, ZYW, DHC, MJT, XMW	
414	and GYS), analyzed the data, and wrote the manuscript. XFL, YZF <sub>2</sub> and QHL conducted air	
415	sampling work and laboratory experiments under the guidance of GHZ, XHB and XMW.	
416	All authors contributed to the refinement of the submitted manuscript.	
417		
418	Acknowledgement	
419	This work was supported by the National Nature Science Foundation of China	
420	(No. 41775124 and 41877307), the National Key Research and Development Program of	
421	China (2017YFC0210104 and 2016YFC0202204), the Science and Technology Project of	
422	Guangzhou, China (No. 201803030032), and the Guangdong Foundation for Program of	

423 Science and Technology Research (No. 2017B030314057).

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## 785 Figure captions

786	Figure 1. Representative mass spectrum for NOCs-containing particles. The ion
787	peaks corresponding to NOCs and oxidized organics are highlighted with red bars.
788	Figure 2. The variation in hourly mean Nfs of the oxidized organics and
789	ammonium that internally mixed with NOCs. Box and whisker plot shows lower,
790	median, and upper lines, denoting the $25^{\text{th}}$ , $50^{\text{th}}$ , and $75^{\text{th}}$ percentiles, respectively; the
791	lower and upper edges denote the 10 <sup>th</sup> and 90 <sup>th</sup> percentiles, respectively.
792	Figure 3. Correlation analysis of (a, c) the RPAs and (b, d) the number of
793	detected NOCs, with the oxidized organics and ammonium in different seasons.
794	Significant ( $p < 0.01$ ) correlations were obtained for both the total observed data and
795	the seasonally separated data. RPA is defined as the fractional peak area of each m/z
796	relative to the sum of peak areas in the mass spectrum and is applied to represent the
797	relative amount of a species on a particle (Jeong et al., 2011; Healy et al., 2013).
798	Figure 4. Comparison between the measured and predicted RPAs for NOCs.
799	Figure 5. (left) PMF-resolved 3-factor source profiles (percentage of total species)
800	and (right) their diurnal variations (arbitrary unit).
801	Figure 6. The dependence of NOCs and the ratio of NOCs to the oxidized organics

802 on RH.

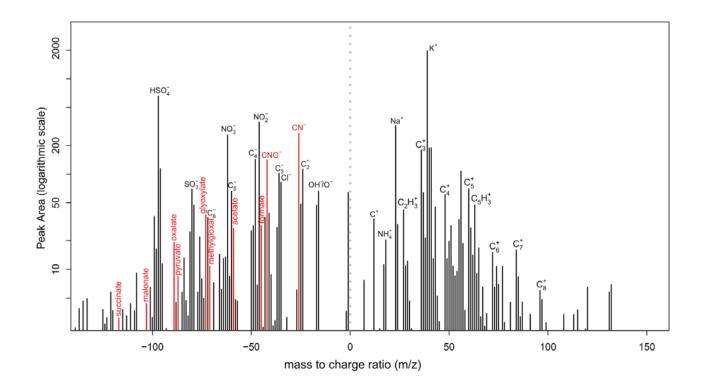
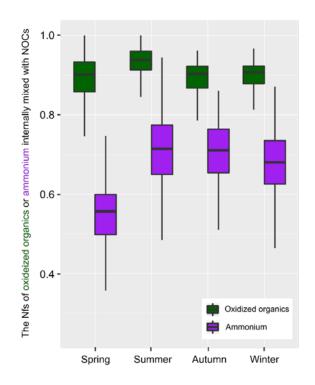


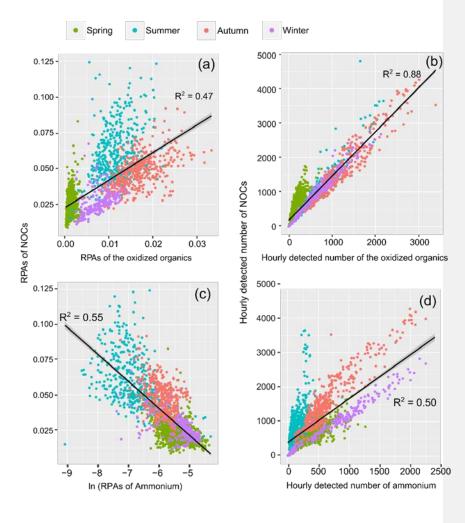


Fig. 1.











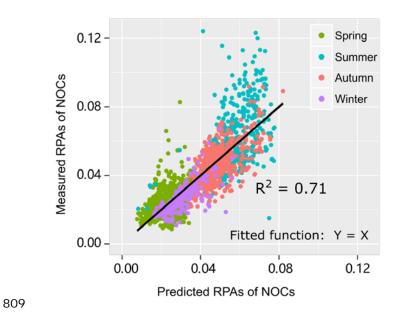
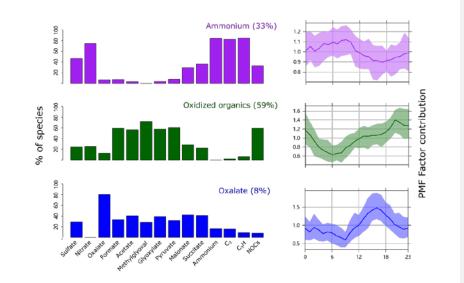




Fig. 4.



813 Fig. 5.

