1	Spatial-temporal Variations and Process Analysis of O₃
2	Pollution in Hangzhou during the G20 Summit
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4	By
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20	Atmospheric Chemistry and Physics
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Abstract

34 Serious urban ozone (O₃) pollution was observed during the campaign of 2016 G20 summit in Hangzhou, China, while other pollutants had been significantly reduced by the short-term emission 35 control measures. To understand the underlying mechanism, the Weather Research Forecast with 36 Chemistry (WRF-Chem) model is used to investigate the spatial and temporal O₃ variations in 37 Hangzhou from August 24 to September 6, 2016. The model is first successfully evaluated and 38 validated for local and regional meteorological and chemical parameters by using the ground and 39 upper-air level observed data. High ozone concentrations, temporarily during most day time of the 40 emission control period and spatially from the surface to the top of the planetary boundary layer, are 41 42 captured in Hangzhou and even the whole YRD region. Various atmospheric processes are further analyzed to determine the influential factors of local ozone formation through the integrated process 43 rate method. Interesting horizontal and vertical advection circulations of O₃ are observed during 44 45 several short periods, and the effects of these processes are nearly cancelled out. As a result, the ozone pollution is mainly attributed to the local photochemical reactions which are not obviously 46 influenced by the emission reduction measures. The ratio of reduction of Volatile Organic 47 Compounds (VOCs) to that of NOx is a critical parameter that needs to be carefully considered for 48 future alleviation of ozone formation. In addition, the vertical diffusion from the upper-air 49 background O₃ also plays an important role in shaping the surface ozone concentration. These 50 results provide insight into urban O₃ formation in Hangzhou and support the Model 51 Intercomparison Study Asia Phase III (MICS-Asia Phase III). 52

53 Keywords: Ozone pollution, WRF-Chem, Spatial-temporal variation, Process analysis, Air quality.

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59 1. Introduction

60 Tropospheric ozone (O_3) is generated by a series of photochemical reactions involving volatile organic compounds (VOCs), nitrogen oxide (NO_x), and carbon monoxide (CO) (Wang et al., 2006). 61 As a primary component of photochemical smog, ground-level O₃ pollution imposes detrimental 62 effects on human health (Ha et al., 2014; Kheirbek et al., 2013) and the ecosystem (Landry et al., 63 64 2013; Teixeira et al., 2011). However, O₃ pollution is a challenging problem worldwide. O₃ levels in 65 cities in the United States and Europe are increasing more than those in the rural areas of these regions, where peak values gradually decreased during 1990-2010 (Paoletti et al., 2014). 66 Nagashima et al., (2017) reported that long-term (1980–2005) trends of increase in surface O₃ over 67 Japan may be primarily attributed to the continental transport that have contributed to 68 photochemical O₃ production. Urban O₃ pollution events have also be observed in developing 69 70 countries, such as Thailand (Zhang and Kim Oanh, 2002) and India (Calfapietra et al., 2016).

Many field monitoring and modeling studies have investigated the photochemical 71 72 characteristics of near-surface O₃ pollution (Tang et al., 2009, 2012; Wang et al., 2013, 2014), the photochemistry of O₃ and its precursors (Xie et al., 2014), the interactions between O₃ and PM_{2.5} 73 (Shi et al., 2015), and the urban O₃ formation (Tie et al., 2013). It is clear that in addition to 74 anthropogenic emissions of O₃ precursors, uncontrollable physical and chemical processes involved 75 in meteorological phenomena significantly modulate changes in O₃ concentration (Xue et al., 2014). 76 In the Yangtze River Delta (YRD) region of China, high O₃ concentrations have been observed 77 (Gao et al., 2016; Jiang et al., 2012). Synoptic patterns related to tropical cyclones may be one 78 79 reason for such high O_3 concentrations (Huang et al., 2005). Jiang et al. (2015) reported that 80 enhanced stratosphere-troposphere exchange (STE) driven by a tropical cyclone abruptly increased O₃ concentrations (21–42 ppb) in the southeast of China during June 12–14, 2014, which has been 81 highlighted as another contributor to near-surface O₃ concentrations under certain conditions (Lin et 82 83 al., 2012, 2015). However, the complex dynamics in atmospheric processes related to O_3 formation 84 are so difficult to identify that the O₃ pollution characteristics and underlying causes have not yet
85 been well understood.

86 Hangzhou, the capital of Zhejiang Province, is located in the center of the Yangtze River Delta which is one of the most developed areas in China. Resultant from local emissions (Wu et al. 2014, 87 Hu et al. 2015) and transboundary transport of aerosol and trace gases transport (Liu et al. 2015; Ni 88 et al. 2018; Zhang et al. 2018), air pollution in Hangzhou has become serious in the recent years. In 89 2016, Hangzhou city would host the 2016 G20 (Group of Twenty Finance Ministers and Central 90 Bank Governors) summit during September 4-6. To improve air quality for this event, 14-day 91 92 temporarily strict air pollution alleviation measures had been taken to reduce air pollutant emissions in Hangzhou and surrounding areas from August 24 to September 6, 2016. The emission control 93 scheme includes a coal-fired power plant capacity 50% reduction since August 24, followed by an 94 95 "odd-even" on-road vehicle restriction since August 28, and further emergent VOC reduction from industrial sectors since September 1 to 6 (Ji et al. 2018; Li et al. 2019; Wu et al. 2019). These short-96 term measures provide a valuable opportunity to investigate the response of air quality to the 97 emission reduction, understand the formation mechanisms of air pollution, and explore effective 98 policies for long-term air pollution control in the local or regional scale. 99

100 The effects of emission control on air pollutants during this G20 Summit have been investigated by several studies using field observations and numerical models. It is demonstrated 101 that almost all major air pollutants including SO₂, NO_x (Li et al. 2019; Wu et al. 2019), fine 102 particles (Ji et al. 2018; Li et al. 2019; Yu et al. 2018; Wu et al. 2019), and VOCs (Zheng et al. 2019) 103 have been significantly reduced during the 14-day control period, except O3. Su et al. (2017) 104 monitored the vertical profiles of ozone concentration in the lower troposphere of Hangzhou during 105 106 the control period by using an ozone lidar. It was found that the ozone concentrations peaked near the top of the planetary boundary layer, and the temporary measures took no immediate effect on 107 108 ozone pollution. Wu et al. (2019) investigated the variation of air pollution in Hangzhou and its surrounding areas during the G20 summit by using monitoring data from five sites, and reported 109

110 that the air quality had been greatly improved by the implementation of the emission control. 111 However, the average O_3 concentration was increased by 19% compared to the same periods of the 112 five preceding years. This unique response of ozone pollution to control measures is not well 113 understood, and of great research interest for better control of ozone pollution in the future.

114 To this end, a regional air quality model, within the framework of the Model Inter-Comparison Study for ASIA phase III (Li et al., 2019), is used to investigate the spatial-temporal characteristics 115 of ozone pollution in Hangzhou during the G20 Summit in the present work. Process analysis is 116 conducted to understand the chemical and physical factors that contribute to O₃ abundance. It is 117 118 found that the serious ozone pollution happened, mainly resultant from the local photochemical reactions which are not under good control by the emission reduction measures. The rest of this 119 paper is organized as follows. Section 2 outlines the methodology and configuration of the model 120 121 system. Section 3 presents the model evaluation, the spatial-temporal characteristics of ozone pollution, and the analysis of related atmospheric processes. Section 4 discusses the underlying 122 causes of O₃ pollution. Finally, a summary is made. 123

124 **2. Methodology**

125 2.1. Regional chemistry modeling system

To investigate the interactions among emissions, meteorological phenomena, and chemical 126 phenomena, the Weather Research Forecast with Chemistry model (WRF-Chem) is used in the 127 present study. The WRF-Chem is a regional online-coupled air quality model that can 128 simultaneously simulate air quality components and meteorological components by using identical 129 130 transport schemes, grid structures, and physical schemes (Grell et al., 2005). Two model domains are designed: an outer domain (horizontal resolution: 30 km) covering East China (20.0°N-44.5°N, 131 99.0°E-126.5°E) and an inner domain (horizontal resolution: 6 km) covering the YRD region 132 (27.6°N-32.7°N, 116.9°E-122.4°E), as shown in Fig.1. The "Lambert conformal conic" projection 133 is applied with the domain center at 34°N, 111°E. A total of 31 vertical layers are used with the 134 model top at 50 hPa. The simulation period is from 17 August to 6 September 2016, and the first-135

136 week simulation is used to spin up the model. Hourly model outputs for 24 August to 06 September
137 are used in the analysis. The gas mechanism CBMZ (Chemical Bond Mechanism Version Z)
138 (Zaveri and Peters, 1999) is used for model simulations. For additional details regarding the model
139 parameterization schemes, please refer to a previous study (Ni et al., 2018).



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141Fig. 1. Double-nested simulation domains. (a) Domain 1: 30 km in East China with 102 (W–E) × 111 (S–N) × 31142(vertical layers) grids; (b) Domain 2: 6 km in the Yangtze River Delta (YRD) region with 100 (W–E) × 115 (S–N)143× 31 (vertical layers) grids. Blue dots denote the air quality monitoring sites. The copyright of the background144map belongs to Google Map © Google.

The meteorological boundary and initial conditions are determined from the global objective final analysis (FNL) data of the National Centers for Environmental Prediction (Kalnay et al., 1996). The FNL data are mapped to domain 1 (East China), and the grid-nudging method (Stauffer et al., 148 1991) is used to reduce the meteorological integral errors. The chemical initial and boundary conditions are dynamically downscaled from the simulation results of model for ozone and related 150 chemical tracers, version 4 (MOZART4) (Emmons et al., 2010).

151 **2.2. Emissions**

152 The 2016 Multiresolution Emission Inventory for China (MEIC, $0.25^{\circ} \times 0.25^{\circ}$; 153 http://www.meicmodel.org/) is used for the outer domain (Fig. 1a) with a spatial resolution of 30 154 km (Li et al., 2017), including species of SO₂, NO_x, CO, NH₃, PM_{2.5}, and VOCs from the power, 155 industrial, residential, transportation, and agricultural sectors. Inventories of finer anthropogenic emissions for the YRD region over the year of 2014 compiled by Shanghai Academy of 156 157 Environmental Sciences are used for the inner domain (Fig. 1b). These inventories have been well 158 documented in previous studies (Huang et al., 2011; Li et al., 2011; Liu et al., 2018). The fine emission inventories include major sectors such as large point sources, industrial sources, mobile 159 sources, and residential sources. The anthropogenic emissions over the YRD region are mainly 160 located over the industrial and urban areas along the Yangtze River as well as over Hangzhou Bay. 161 In this study, the emission inventories for the two domains are projected into horizontal and vertical 162 163 grids as hourly emissions, with temporal and vertical profiles obtained from Wang et al. (2011). VOCs emissions are categorized into modeled species, according to von Schneidemesser et al. 164 (2016). In addition, biogenic emissions are generated offline using the Model of Emission of Gases 165 166 and Aerosols from Nature (MEGAN) (Guenther et al., 2006). Dust emissions are calculated online from surface features and meteorological fields by using the Air Force Weather Agency and 167 Atmospheric and Environmental Research scheme (Jones et al., 2011). Other emissions, such as 168 those from biomass burning, aviation, and sailing ships, accounting for very small fraction during 169 170 this period, are therefore not considered here. However, it is worth noting that these base inventories have been modified in the simulation to reflect the realistic emissions according to the 171 control measures taken in the period presented in the introduction. 172

173 2.3. Atmospheric processes analysis

To understand the underlying mechanism of O₃ formation, individual physical and chemical processes of O₃ formation are investigated by using the integrated process rate (IPR) analysis in the WRF-Chem model (Jfffries and Tonnesen, 1994). The IPR analysis differentiates changes in pollutant concentrations from individual atmospheric process which quantitatively elucidates the contributions of each process, mainly including advection, diffusion, emission, deposition, clouds process, aerosol and gaseous chemistry. The IPR analysis has been widely applied and demonstrated to be an effective tool for investigating the relative importance of individual processes and 181 interpreting O_3 concentrations (Goncalves et al., 2009; Tang et al., 2017; Shu et al., 2016). In the 182 present work, we consider gas chemistry, vertical diffusion, horizontal and vertical advections as the 183 main atmospheric processes for O_3 formation. Other processes, such as cloud process and horizontal 184 diffusion, play minor roles and are thus not considered.

185 2.4. Evaluation metrics

186 To increase the confidence in interpretations of model results, model outputs should first be evaluated based on observations. Accordingly, the model results derived from domain 2 are 187 compared with hourly surface observational data obtained from 96 air quality monitoring sites in 188 189 the YRD region (blue dots, Fig. 1b) in this study. These observational data are downloaded from http://www.pm25.in, and O₃ as well as its precursor NO₂ are evaluated, in terms of statistical 190 measures, namely the mean fractional bias (MFB), the mean fractional error (MFE), and the 191 192 correlation coefficient (R), following the recommendation of the US Environmental Protection Agency (US EPA, 2007). Additionally, the meteorological parameters are evaluated based on the 193 194 observational data, including temperature at 2 m (T2), relative humidity at 2 m (RH2), 10 m wind speed (WS10) and direction (WD10), from the Meteorological Assimilation Data Ingest System 195

196 Table 1. Discrete statistical indicators used in the model evaluation

Metrics	Definition	Range
Mean Fractional Bias (MFB)	$MFB = \frac{2}{N} \sum_{i=1}^{N} \frac{S_i - O_i}{S_i + O_i} \times 100\%$	-200% to 200%
Mean Fractional Error (MFE)	$MFE = \frac{2}{N} \sum_{i=1}^{N} \frac{ S_i - O_i }{S_i + O_i} \times 100\%$	0 to 200%
Correlation Coefficient (r)	$r = \frac{\sum_{i=1}^{N} (S_i - \overline{S})(O_i - \overline{O})}{\sqrt{\sum_{i=1}^{N} (S_i - \overline{S})^2 \sum_{i=1}^{N} (O_i - \overline{O})^2}}$	0 to 1
Mean Bias (MB)	$MB = \frac{1}{N} \sum_{i=1}^{N} (S_i - O_i)$	$-\infty$ to $+\infty$
Gross Error (GE)	$GE = \frac{1}{N} \sum_{i=1}^{N} \left S_i - O_i \right $	0 to $+\infty$
Root Mean Square Error (RMSE)	$RMSE = \sqrt{\frac{1}{N} \sum_{i=1}^{N} (S_i - O_i)^2}$	0 to $+\infty$

N is the number of samples. S_i and O_i are values of simulations and observations at time or location *i*, respectively.

197 (https://madis.noaa.gov). Following the study of Zhang et al. (2014), commonly used mean bias 198 (MB), gross error (GE), and root mean square error (RMSE) are calculated as the statistical 199 indicators. All used statistical indicators are summarized in Table 1.

Besides the above evaluation of single-point based time series results, the vertical spatial distribution of modeled O₃ in Hangzhou is also evaluated by comparisons with observed differential absorption LiDAR (DIAL) data (Su et al. 2017). In the DIAL technique, the mean gas concentration over a certain range interval is determined by analyzing the LiDAR backscatter signals for laser wavelengths tuned "on" (λ_{on}) and "off" (λ_{off}) in a molecular absorption peak of the gas under



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Fig. 2. Comparison of modeled air pollutant concentrations against measurements in 96 monitoring sites over VRD region during August 24–September 6, 2016: Mean fractional bias (MFB), mean fractional error (MFE) and Pearson's correlation coefficient (r) of O₃ (a–c) and NO₂ (d–f), respectively.

209 investigation (Browell et al., 1998). In the DIAL data of O_3 , the vertical height available is from 0.3 210 km to 3 km due to the limitations of the signal-to-noise ratio and detection range.

211 **3. Results**

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212 **3.1. Model performance**

We first evaluate the overall performance of WRF-Chem for the YRD region by incorporating data from the 96 air quality monitoring sites. Specifically, the maximum daily 8 h (MDA8) ozone and daily mean NO₂ concentrations at the surface are used. The spatial distributions of MFB and MFE for O_3 and NO_2 are illustrated in Fig. 2. In general, the model-simulated air pollutant concentrations agree well with the observations, with MFB and MFE for most of the sites meeting



Fig. 3. Comparisons of modeled and observed concentrations of the air pollutants from 96 air quality monitoring sites across the YRD from 24 August to 06 September 2016 (1,344 pairs). Scatter plots for MFB and MFE of (a) O_3 and (b) NO₂. Performance goals (red box) for O_3 are the benchmarks. Scatter plots for daily observed and modeled (c) O_3 and (d) NO₂.

the benchmarks (MFB<15%; MFE<35%) (US EPA, 2007). A scatter plot of MFB/MFE is shown in Fig.3, further demonstrating the capability of the present model in reproducing the observations, which is also supported by the high correlation between model and observation (Fig. 3c, d).</p>



Fig. 4. Modeled air pollutants and meteorological parameters compared with measurements at the Hangzhou monitoring site from 24 August to 06 September 2016. Surface concentrations of (a) O_3 and (b) NO_2 , (c) temperature at 2 m (T2), (d) relative humidity at 2 m (RH2), (e) wind speed at 10 m (WS10) and (f) wind direction at 10 m (WD10).

231 After the above overall evaluation of the present model in the whole YRD region, the site of Hangzhou will be focused on for further analysis. The time series of hourly simulated and observed 232 air pollutants (O₃, Fig. 4a; NO₂, Fig. 4b) and meteorological factors (T2, Fig. 4c; RH2, Fig. 4d; 233 234 WS10, Fig. 4e; and WD10, Fig. 4f) at Hangzhou are presented in Fig. 4. It is found that all modeled 235 data are statistically significantly correlated with the observed data at the 95% level. The MFB and MFE for both O₃ and NO₂ are well below the benchmarks (MFB/MFE: 15%/35%; US EPA, 2007) 236 237 and the observed diurnal variations are well reproduced. For meteorological parameters, $\leq \pm 0.5^{\circ}$ C, $GE \le 2.0$ °C), 10 m wind speed (MB ≤ 0.5 m/s, RMSE ≤ 2.0 m/s) and 10 m wind direction (MB \le 238 $\pm 10^{\circ}$, GE $\leq 30^{\circ}$). McNally (2009) suggested a relaxed benchmark for 2 m temperature (MB \leq 239 $\pm 1.0^{\circ}$ C). In this study, the 10 m wind speed and wind direction (Fig. 3e, f) results are well within 240

the benchmarks. The GE of 2 m air temperature (1.9° C; Fig. 3c) also satisfies the criteria, but the MB is slightly higher (-1.6° C) which has also been noted in a previous study (Zhang et al., 2014). These comparisons further demonstrate that the present model is able to correctly predict the time series of both meteorological parameters and air pollutants of O₃ and NO₂ in Hangzhou.



Fig. 5. Vertical comparison of hourly (a) observed (from differential absorption LiDAR) and (b) simulated O₃ concentrations (ppb) in Hangzhou from 24 August to 06 September 2016. Purple regions in the top panel denote invalid data with a low signal-to-noise ratio. To facilitate direct comparison, the red dashed line is added to indicate the ozone level recorded for the same time periods (starting from 12:00, 24 August) and vertical heights (0.3–3 km) in the observations and simulation results.

To further evaluate the capability of the model to predict the vertical structure of ozone concentration, the vertical distribution of the modeled O_3 in Hangzhou from 24 August to 06 September 2016 is qualitatively compared with the DIAL data, as shown in Fig. 5. It is interesting to find that the present model can successfully predict the spatial-temporal distribution of ozone in Hangzhou. All observed major features of ozone are well captured by the model. This gives us high confidence and lays a solid foundation for further exploring the pollution characteristics and influencing factors of ozone in Hangzhou during the G20 summit.

258 **3.2.** Spatial-temporal variations of O₃ pollution

To discuss spatial-temporal characteristics of O_3 pollution in Hangzhou, the whole emission control period can be divided into three stages according to the reduction intensity of the measures. August 24-27, 2016 is the first stage (S1) during which industrial and construction emission controls were implemented. During the second stage (S2, August 28-31), traffic restrictions were 263 further added. September 1-6 2016 is the third stage (S3) with the emergent VOCs control further implemented. Figs. 4(a) and 4(b) in the above section also present the temporal evolution of O₃ and 264 265 its precursor NO₂ in Hangzhou city during the emission control period of G20 summit. It is evident 266 that the NO₂ has been significantly reduced by the emission control measures and the concentration is well below the national level-II standard of 200 μ g/m³. However, the concentration of O₃ keeps 267 high levels for the whole 14 days, with 7 days of MDA8 are above and 4 days are close to the 268 national level-II standard (GB-3095-2012) of 160µg/m³. This serious O₃ pollution indicates that the 269 emission control measures seem to make no obvious effect on ozone, which is consistent with 270 271 previous observations (Su et al. 2017; Wu et al. 2019). The diurnal variation of O₃ is similar for the





Fig. 6. Synoptic circulation in East Asia during the 2016 G20 summit. Weather charts for four representative
periods at 08:00 LST on (a) 25 August, (b) 27 August, (c) 31 August, and (d) 06 September 2016. LST: Local
Sidereal Time; H: High-pressure system; L: Low-pressure system. The red triangle denotes the location of
Hangzhou city.

three stages with a peak value at the time around 16:00 and a valley value at the time around 8:00 of each day. However, the variation magnitude in Stage 2 is obviously lower than those of other stages, which will be further discussed later.

Fig. 5 also clearly shows this diurnal variation of O_3 in the ground level. However, nocturnal O₃-rich mass is observed during certain periods in the upper air (approximately 1 km), such as August 25, August 31, and September 3, which makes an n-shaped distribution pattern of the O₃ in the upper air. This kind of spatial distribution of ozone will promote vertical exchange of O₃ in the area. In general, high concentrations of O₃ appear vertically until the top of the planetary boundary layer (PBL, approximately <2 km), suggesting the ozone pollution is not a local but a regional phenomenon in the whole low-level (from surface to close to the PBL height) region.

287 Considering the synoptic circulation is closely related to regional O₃ abundance, four 288 representative surface weather charts obtained from the Korea Meteorological Administration are presented in Fig. 6. In the early stage, strong and uniform high-pressure fields covered vast regions 289 of southeastern China, and a tropical cyclone moved northeastward over the East China Sea (Fig. 290 6a). In the middle stage (Fig. 6b), the tropical cyclone approached the YRD region, bringing strong 291 292 north wind fields to this area. As a result, the long narrow rain band arrived in Hangzhou (red triangle) on 27 August 2016. In the later stage (Fig. 6c), the cyclone continuously moved and 293 eventually hit the land, and the tropical high in the YRD region recovered gradually. Finally, the 294 295 cyclone faded, and a rainstorm appeared over most of the YRD region (Fig. 6d).

The typical hourly vertical and horizontal O_3 distributions in the YRD region are further presented in Fig.7. The wind fields are also included for better understanding. For stagnation days with weak wind fields and strong radiation before or after the tropical cyclone, meteorological conditions are unfavorable for pollutant dispersion. As a result, O_3 pollution is more regional and intense, with an hourly peak O_3 concentration of 250 µg/m³ appeared within the planetary boundary layer in the whole YRD region, as shown in Figs. 7a and 7c. In these conditions, photochemical



302 303 Fig. 7. Surface and low-level O₃ distributions (μ g m⁻³) and wind fields (vectors, m s⁻¹) for representative episodes. 304 (a) Stagnant weather before the tropical cyclone, (b) pollutant transport when the tropical cyclone approached, and 305 (c) stagnant weather after the cyclone. The red line denotes the cross-section line of low-level O₃ distributions. 306 The red triangle denotes the location of Hangzhou.

307 reactions dominate the ozone formation and accumulation. This phenomenon is consistent with the 308 satellite-derived tropospheric O_3 distribution in the area (Su et al. 2017), and is also supported by 309 the observed ozone data from the 96 sites in the YRD region as shown in Fig. 3c. During the 14-day 310 emission control period of G20 summit, 52% of the observed ozone samples from the 96 sites are above the China's national level-II standard ($160\mu g/m^3$), suggesting that regional ozone pollution 311 appears in the YRD region during the study period. As the cyclone approached on 27 August, a 312 313 large belt of O_3 mass appeared in the upwind direction and moved toward Hangzhou under a 314 prevailing north wind field (Fig. 7b). Regional pollutant transport may play an important role under 315 this condition. However, because of the rain and cooling effects from the cyclone, the ozone 316 concentration is relatively low in the whole YRD region.



317 3.3. Process analysis of O₃ formation

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321 To further investigate the underlying mechanism of O_3 pollution, hourly variations in the 322 change rate of low-level O_3 resulting from different physical and chemical processes are presented 323 in Fig.8. It is evident that gas chemistry is the dominant factor to the strong generation of abundant 324 O_3 in the entire planetary boundary layer (<2 km) in the daytime but causes small amount of 325 depletion of O₃ at the near-surface height (<0.3 km) in the nighttime (Fig. 8a). High concentration 326 of O_3 diffuses from the upper layer downward to the ground through vertical diffusion during the 327 whole study period, which is obvious in the daytime (Fig. 8b). However, this effect is relatively 328 weak compared to other processes. Horizontal and vertical advections seem to play more important 329 roles in shaping the near-surface O_3 , as indicated in Figs. 8c and 8d. Several interesting dynamic O_3 circulations are observed between the near-surface and upper-air regions and indicated by the 330 331 dashed boxes. During the periods of 24-26 August and 31 August to 2 September 2016, O₃-rich 332 mass in the lower layer (<1 km) travels to Hangzhou through horizontal advection, and is then 333 transported upward to the higher layer through vertical advection. At this higher layer, the mass



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335 Fig. 9. (a) Daytime mean (08:00–17:00 LST) change rate of simulated surface O_3 (ppb h⁻¹; left *y* axis) resulted 336 from gas chemistry, vertical diffusion, horizontal and vertical advections in Hangzhou. (b) Correlation of daytime 337 mean gas chemistry generation (ppb h⁻¹; left *y* axis) and observed surface-level maximum for daily 8h 338 concentration of O_3 (ppb; right *y* axis) in Hangzhou. China's national level-II standard is approximately 75 ppb 339 (160 µg m⁻³).

340 subsequently travels away from Hangzhou to other places through horizontal advection in a circular 341 manner. This phenomenon might be associated with the urban heat island circulation (Lai and 342 Cheng, 2009). However, during the period of September 3-6 2016, similar circulation phenomenon 343 is observed, but the flow direction is reverse. The O3-rich mass travels downward to the ground through vertical advection and is then transported to surrounding regions through horizontal 344 advection. This downward circulation is also related to the meteorological conditions after the 345 346 cyclone. In addition, the horizontal and vertical advections of O_3 take on a chaotic status during 347 August 27-30, 2016, suggesting that complicated variable meteorological conditions happened in 348 the time. This is also the reason for the lower magnitude of diurnal variation in Stage 2.

Fig. 9 shows the daytime mean change rate of simulated O₃ at the ground level resulted from 349 various atmospheric processes and the correlation of gas chemistry generation and observed 350 351 maximum for daily 8h concentration of O₃. As a whole, the main sources of local surface ozone in Hangzhou are from gas chemistry, vertical diffusion, and horizontal advection, with mean 352 production rates of 1.9, 3.3, and 6.7 ppb h⁻¹, respectively, from August 24 to September 6, 2016, 353 and the major sink is vertical advection. However, during some days such as September 5-6, the gas 354 chemistry consumes O₃ while the vertical advection increases it. In general, strong net horizontal 355 356 and vertical advections of O3 are observed for most days of the period, except August 27-28 during



Fig. 10. Simulated hourly downward short wave flux at ground surface in Hangzhou (W m⁻²) during August 24 to
September 6, 2016.

360 which the strongest northwest cold winds (Fig. 4e) occurred and made the net advections of O_3 361 negligible. Similar to Fig. 8, dynamic O_3 circulations are observed for the periods of August 24-26, 362 August 31 to September 2, and September 5-6. Particularly, the circular direction is reverse during 363 September 5-6 and the net gas chemistry is to consume ozone due to weak solar radiation in the 364 days as shown in Fig. 10.

365 In addition, the variation trend of the daytime mean production rate of gas chemistry is consistent with the observed MDA8 concentration and the local chemical generation has large 366 positive correlation (Pearson's r = 0.77) with the observed MDA8 concentrations (Fig. 8b). This 367 368 indicates a trade-off effect among vertical diffusion, horizontal advection, and vertical advection. High O₃ concentrations (i.e., 25 August 2016 MDA8: 98 ppb) are always accompanied by strong 369 radiation and prolific generation of gas chemical reactions. It is also interesting to find that vertical 370 371 diffusion may partially compensate for gas chemistry when the chemical reaction rate is relatively low or negative. For example, during August 26-27 and September 5-6, the vertical diffusion rates 372 are higher than the chemical production rates. The low O₃ episode on these periods may result from 373 local chemical consumption. 374

375 4. Discussion

376 The above results demonstrate that high ozone concentrations are observed, temporarily during most day time of the emission control period of G20 summit, and spatially in Hangzhou and even 377 the whole YRD region, from the surface to the top of the planetary boundary layer. Strong 378 horizontal and vertical advections appear, but they form circulations due to special meteorological 379 380 conditions so that the effects of them almost cancel each other out. As a result, the serious ozone 381 pollution in Hangzhou is mainly resultant from the local photochemical reactions. When the photochemical reactions are weak, the vertical diffusion from the upper-air notable background O₃ 382 further compensates for the local surface ozone concentration. Therefore, it is of great importance to 383 384 understand why the strict emission control measures make no obvious effect on the local photochemical reactions of ozone generation. 385

386 Chemical generation of O_3 is the net effect of photochemical generation and titration consumption. VOC oxidation (Jenkin et al., 1997; Sillman, 1999) in photochemical reactions 387 provides critical oxidants (i.e., RO₂) that efficiently convert NO to NO₂, resulting in further 388 389 accumulation of O₃ (Wang et al., 2017). The chemical generation of O₃ is controlled by NO_x and VOCs depending on which substance is lack in the reactions. As a consequence, there are two 390 sensitivity regimes of O₃ production, namely, the NO_x-limited and VOC-limited regimes. Previous 391 studies have shown that the sensitivity pattern of surface O₃ formation in Hangzhou is dominant by 392 the VOCs-limited regime (Yan et al. 2016; Li et al., 2017; Su et al., 2017). In this regime, if the 393 394 regional reduction of VOCs is much higher than that of NO_x, the O₃ concentration can be reduced, otherwise if the regional reduction of VOCs is much less than that of NO_x, the inhibitory effect of 395 NO_x on O₃ generation will be weakened, and the O₃ concentration will increase remarkably (Wang 396 397 et al. 2015). According to the studies of Su et al. (2017), Zheng et al. (2019), and Wu et al. (2019), it can be deduced that NO_x has been significantly reduced by about 60%, at least two times of the 398 reduction of VOCs in Hangzhou. The influence of stringent emission control measures on VOCs is 399 not as immediate and effect as that on NO_x, which is associated with the fact that there was a large 400 amount of biogenic VOC emission in Hangzhou and surrounding regions (Liu et al. 2018; Wu et al. 401 2020). In fact, the average temperature during the study period is as high as around 31°C (Fig. 4c), 402 which facilitates the biogenic VOC emissions and photochemical reactions. As a result, the 403 photochemical generation of O₃ was not under control and high concentration of ozone appeared. 404 However, it is worth noting that after the emergent VOCs control measures had been implemented 405 in the area during the third stage, the net generation rate of O₃ gradually reduces since September 2, 406 2016, leading to a period of relatively low ozone concentration together with other meteorological 407 408 effects. These discussions implicate that to alleviate ozone pollution, the ratio of reduction of VOCs to that of NO_x is the key parameter based on the O_3 -NO_x-VOCs sensitivity analysis. As the biogenic 409 410 VOCs are important sources of the total VOCs in the YRD region, it is necessary to balance the 411 reduction of NO_x to make the ratio within effective regime in the future.

412 **5.** Conclusions

413 To understand the unique response of ozone to short-term emission control measures during the G20 summit in Hangzhou, the spatial-temporal characteristics and process analysis of O_3 414 pollution are investigated by using the WRF-Chem model. Statistical evaluations of meteorological 415 416 and chemical parameters suggest that the model system is able to reasonably predict the observed data for both the ground and upper-air levels in MICS-Asia III. High ozone concentrations are 417 418 observed, temporarily during most day time of the emission control period of G20 summit, and spatially in Hangzhou and even the whole YRD region, from the surface to the top of the planetary 419 boundary layer. Horizontal and vertical advection circulations are captured in Hangzhou, with 420 horizontal advection the source and vertical advection the sink of the surface O_3 in Hangzhou. 421 Consequently, the serious ozone pollution is mainly resultant from the local photochemical 422 reactions which are not under good control by the emission reduction measures. As the surface O₃ 423 formation in Hangzhou is dominant by the VOCs-limited regime, the significant reduction of NO_x 424 425 compared to that of VOCs is unfavorable to chemical generation of O₃. The ratio of reduction of VOCs to that of NO_x based on the O₃-NO_x-VOCs sensitivity analysis is a critical parameter for 426 reduction of ozone formation from photochemical reactions. In addition, it is found that the vertical 427 428 diffusion from the upper-air notable background O₃ also plays an important role in shaping the 429 surface ozone concentration when the photochemical reactions are weak.

430 Author contribution

431 Zhizhen Ni: Data curation, Investigation, Writing - original draft. Kun Luo: Methodology,
432 Resources, Writing - review & editing, Supervision. Yang Gao: Formal analysis, Methodology,
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437 Acknowledgments

This work was financially supported by special funds from the Ministry of Environmental Protection of China (No. 201409008-4) and the Zhejiang Provincial Key Science and Technology Project for Social Development (No. 2014C03025). We would like to thank the US National Oceanic and Atmospheric Administration for its technical support in WRF-Chem modeling. Highresolution emission inventories were provided by the Institute of Environmental Science, Shanghai, China, and the official documents of emission control policies were obtained from the Hangzhou Environmental Monitoring Center.

445 Competing interests

446 The authors declare that they have no conflict of interest.

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