Responses to comments by editor and reviewer #2

Comments by editor

Both referees have evaluated your revised manuscript. The referee #1 provided several suggestions for a further minor revision. Please see the report. Regarding the comment on Figure 4, I suggest defining the abbreviations "AN_" and "BB_" in the figure caption, as the regions have been named in Figure 1.

Response: We appreciate the suggestion. The abbreviations are defined as suggested, i. e., Lines 772-773 "Figure 4. Contributions of anthropogenic sources (prefixed "AN_" in the legend) and open biomass burning ("BB_")…".

Comments by reviewer #2

The authors well solved my concerns in the revised manuscript. I only have few minor comments now. Please see below:

Line 92: The observed BC concentration could be biased by a factor of 2 in Sharma et al. (2017).

Response: In light of the point, the sentence is revised to reflect the larger bias of 2 in Lines 92-93 to "The observations of BC that are used for model comparisons may be biased by a factor of 2 depending on the method used (Sinha et al., 2017; Sharma et al., 2017)".

Line 132: Please define 5000 m as "high altitudes" here.

Response: Revised in Line 133-134 from "…5000 m (4750–5250 m), were simulated" to "… high altitudes (4750–5250 m), were simulated". Relevant corrections were done through the context to avoid iteration.

Figure 4: Please present the full name before using the abbreviations (e.g., AN_EUR).

Response: The abbreviations are added in the figure caption in Lines 772-773 as "Figure 4. Contributions of anthropogenic sources (prefixed "AN_" in the legend) and open biomass burning ("BB_")…".

Line 48: Many recent studies examined the impact of BC on Arctic climate (e.g., Sand et al., 2016; Yang et al., 2019). In addition to BC, tropospheric ozone, methane, other short-lived climate forcers, like sulfate aerosol (Yang et al., 2018), could also affect the Arctic climate.

Response: These recent progresses were added in L48-49 as "Short-lived climate pollutants such as black carbon (BC) particles (e.g., Sand et al., 2016; Yang et al., 2019), sulfate aerosol (Yang et al., 2018), tropospheric ozone, …", and corresponding references added.

Technical correction:

Line 140, "life time" is corrected to "lifetime".

Flexpart v10.1 simulation of source contributions to Arctic black carbon

Chunmao Zhu[1], Yugo Kanaya[1,2], Masayuki Takigawa[1,2], Kohei Ikeda[3], Hiroshi Tanimoto[3], Fumikazu Taketani[1,2], Takuma Miyakawa[1,2], Hideki Kobayashi[1,2], Ignacio Pisso[4]

[1]Research Institute for Global Change, Japan Agency for Marine–Earth Science and Technology (JAMSTEC), Yokohama 2360001, Japan

[2]Institute of Arctic Climate and Environmental Research, Japan Agency for Marine–Earth Science and Technology, Yokohama 2360001, Japan

[3]National Institute for Environmental Studies, Tsukuba 305-8506, Japan

[4]NILU – Norwegian Institute for Air Research, Kjeller 2027, Norway

Correspondence to Chunmao Zhu (chmzhu@jamstec.go.jp)

Abstract

The Arctic environment is undergoing rapid changes such as faster warming than the global average and exceptional melting of glaciers in Greenland. Black carbon (BC) particles, which are a short-lived climate pollutant, are one cause of Arctic warming and glacier melting. However, the sources of BC particles are still uncertain. We simulated the potential emission sensitivity of atmospheric BC present over the Arctic (north of 66° N) using the Flexpart Lagrangian transport model (version 10.1). This version includes a new aerosol wet removal scheme, which better represents particle-scavenging processes than older versions did. Arctic BC at the surface (0–500 m) and high altitudes (4750–5250 m) is sensitive to emissions in high latitude (north of 60° N) and mid-latitude (30–60° N) regions, respectively. Geospatial sources of Arctic BC were quantified, with a focus on emissions from anthropogenic activities (including domestic biofuel burning) and open biomass burning (including agricultural burning in the open field) in 2010. We found that anthropogenic sources contributed 82 % and 83 % of annual Arctic BC at the surface and high altitudes, respectively. Arctic surface BC comes predominantly from anthropogenic emissions in Russia (56 %), with gas flaring from the Yamalo-Nenets Autonomous Okrug and Komi Republic being the main source (31 % of Arctic surface BC). These results highlight the need for regulations to control BC emissions from gas flaring to mitigate the rapid changes in the Arctic environment. In summer, combined open biomass burning in Siberia, Alaska, and Canada contributes 56–85 % (75 % on average) and 40–72 % (57 %) of Arctic BC at the surface and high altitudes, respectively. A large fraction (40 %) of BC in the Arctic at high altitudes comes from anthropogenic emissions in East Asia, which suggests that the rapidly growing economies of developing countries could have a non-negligible effect on the Arctic. To our knowledge, this is the first year-round evaluation of Arctic BC sources that has been

performed using the new wet deposition scheme in Flexpart. The study provides a scientific basis for actions to mitigate the rapidly changing Arctic environment.

1 Introduction

The Arctic region has experienced warming at a rate twice that of the global average in recent decades (Cohen et al., 2014). The Arctic cryosphere has been undergoing unprecedented changes since the mid-1800s (Trusel et al., 2018). Glacier cover in Greenland reached its historically lowest level in summer 2012 (Tilling et al., 2015). Evidence indicates that the emissions and transport of greenhouse gases and aerosols to the Arctic region are contributing to such warming and melting of snow and ice (Keegan et al., 2014; Najafi et al., 2015). Short-lived climate pollutants such as black carbon (BC) particles (e.g., Sand et al., 2016; Yang et al., 2019), sulfate aerosol (Yang et al., 2018), tropospheric ozone, and methane greatly affect the Arctic climate (AMAP, 2015; Quinn et al., 2008).

BC particles are emitted during incomplete combustion of fossil fuels, biofuels, and biomass. BC warms the atmosphere by direct absorption of solar radiation. The deposition of BC on snow and ice surfaces accelerates their melting through decreasing albedo, which contributes to the rapid loss of glaciers. In the Arctic region, ground-based observations have indicated that BC shows clear seasonal variations, with elevated mass concentrations in winter and spring (the so-called Arctic haze) and low values in summer (Law and Stohl, 2007). Such seasonal variations are explained by increased transport from lower latitudes in the cold season and increased wet scavenging in the warm season (Shaw, 1995; Garrett et al., 2011; Shen et al., 2017).

The presence of BC particles in the Arctic is mainly attributed to emissions in high-latitude regions outside the Arctic, such as northern Europe and Russia (Stohl, 2006; Brock et al., 2011). This is partly caused by the polar dome (Stohl, 2006), which is formed because of the presence of constant potential temperature near the surface. The emissions in high-latitude regions are transported to the Arctic region and trapped in the dome, which increases the

surface concentration. Recently, Schmale et al. (2018) suggested that local emissions from within the Arctic are another important source, and these are expected to increase in the future.

Although numerous studies have been performed, results regarding regional contributions of BC sources in the Arctic are still inconclusive. For example, ground-based observations and Lagrangian transport model results reported by Winiger et al. (2016) showed that BC in Arctic Scandinavia is predominantly linked to emissions in Europe. Over the whole Arctic region (north of 66° N), Russia contributes 62 % to surface BC in terms of the annual mean (Ikeda et al., 2017). Gas flaring in Russia has been identified as a major (42 %) source of BC at the Arctic surface (Stohl et al., 2013). Xu et al. (2017) found that anthropogenic emissions from northern Asia contribute 40–45 % of Arctic surface BC in winter and spring. However, the results of some other studies have suggested that Russia, Europe, and South Asia each contribute 20–25 % of BC to the low-altitude springtime Arctic haze (Koch and Hansen, 2005). Sand et al. (2016) found that the surface temperature in the Arctic is most sensitive to emissions in Arctic countries, and Asian countries contribute greatly to Arctic warming because of the large absolute amount of emissions. With these large disagreements among studies, it is thus necessary to unveil BC sources in the Arctic with high precision simulations.

Various models have been used to investigate BC sources in the Arctic. Depending on the simulation method, these models are generally categorized as Lagrangian transport models (Hirdman et al., 2010; Liu et al., 2015; Stohl et al., 2006, 2013), chemical transport models (Ikeda et al., 2017; Koch and Hansen, 2005; Qi et al., 2017; Shindell et al., 2008; Wang et al., 2011; Xu et al., 2017), and global climate models (Ma et al., 2013; Schacht et al., 2019; H. Wang et al., 2014) (Table 1). The treatment of wet-scavenging parameterizations is a key

factor affecting the model performance, which determines the uncertainties related to BC particle removal (Kipling et al., 2013; Schacht et al., 2019; Q. Wang et al., 2014). The use of emission inventories is another important factor that affects the simulation results (Dong et al., 2019). The observations of BC that are used for model comparisons may be biased by a factor of 2 depending on the method used (Sinha et al., 2017; Sharma et al., 2017). There are still large uncertainties regarding the sources of BC in the Arctic with respect to emission sectors (anthropogenic sources and open biomass burning) and geospatial contributions (Eckhardt et al., 2015).

Deleted: 30 %

The FLEXible PARTicle dispersion model (Flexpart) had been used to investigate the transport pathways and source contributions of BC in the Arctic (Stohl et al., 1998, 2006, 2013). Of Flexpart model up to version 9, wet removal was treated considering below-cloud and within-cloud scavenging processes (Hertel et al., 1995; McMahon and Denison, 1979), which depends on cloud liquid water content, precipitation rate and the depth of the cloud. However, clouds were parameterized based on relative humidity, clouds frequently extended to the surface and at times no clouds could be found in grid cells, with unrealistic precipitation (Grythe et al., 2017). Recently, version 10 of Flexpart had been developed in which cloud is differentiated into liquid, solid, and mixed phase, the cloud distribution is more consistent with the precipitation data (Grythe et al., 2017). This improvement in the cloud distribution and phase leads to a more realistic distribution of below-cloud and in-cloud scavenging events. In this study, we quantified region-separated sources of BC in the Arctic in 2010 by using Flexpart v10.1. We first evaluated the model performance by comparing the results with those based on observations at surface sites. The source contributions of emission sectors and geospatial contributions were evaluated by incorporating the Arctic BC footprint into the emission inventories.

2 Materials and methods

2.1 Transport model

The Flexpart model (version 10.1) was run in backward mode to simulate BC footprints in the Arctic region. The calculation of wet deposition was improved compared with those in previous versions because in-cloud scavenging and below-cloud scavenging of particles were separately calculated (Grythe et al., 2017). In previous versions of Flexpart, in the in-cloud scavenging scheme, the aerosol scavenging coefficient depended on the cloud water content, which was calculated according to an empirical relationship with precipitation rate, in which all aerosols had the same nucleation efficiency (Hertel et al., 1995 ; Stohl et al., 2005). In the new version, the in-cloud scavenging scheme depends on the cloud water phase (liquid, ice, or mixed phase). Aerosols were set as ice nuclei for ice clouds and as cloud condensation nuclei for liquid-water clouds, respectively. For mixed-phase clouds, it was assumed that 10 % of aerosols are ice nuclei and 90 % are cloud condensation nuclei, because BC is much more efficiently removed in liquid water clouds than in ice clouds (Cozic et al., 2007; Grythe et al., 2017). The below-cloud scavenging scheme can parameterize below-cloud removal as a function of aerosol particle size, and precipitation type (snow or rain) and intensity. The biases produced in simulations using the new scheme are therefore smaller than those in the old scheme for wet deposition of aerosols, especially at high latitudes (Grythe et al., 2017).

The Arctic region is defined as areas north of 66° N. The potential BC emission sensitivities at two heights in the Arctic region, i.e., the surface (0–500 m) and high altitudes (4750–5250 m), were simulated. The Flexpart outputs were set as gridded retention times. We performed tests at 500, 2000, and 5000 m, and chose 500 m as the upper boundary height of the model output. The model was driven with operational analytical data from the

Deleted: 5000 m

European Centre for Medium-Range Weather Forecasts (ECMWF) at a spatial resolution of 1° × 1° with 61 vertical levels. Temporally, ECMWF has a resolution of 3 h, with 6 h analysis and 3 h forecast time steps. The simulation period was set at 60 days backward starting from each month in 2010. The maximum lifetime of BC was set at 20 days because its suspension time in the upper atmosphere during long-range transport is longer than that at the surface level (Stohl et al., 2013). We implemented the wet deposition scheme in the backward calculations, but it was not represented in the default setting (Flexpart v10.1, https://www.flexpart.eu/downloads, obtained 10 April 2017).

Deleted: life time

The chemistry and microphysics could not be resolved by Flexpart. The model therefore ignores hydrophobic to hydrophilic state changes and size changes of BC, and assumes that all BC particles are aged hydrophilic particles. This may lead to an overestimation of BC removal and hence force underestimation of simulated BC concentration, especially of fossil fuel combustion sources where BC could be in the hydrophobic state for a few days. A logarithmic size distribution of BC with a mean diameter of 0.16 μm and a standard deviation of 1.96, in accordance with our ship observations in the Arctic, was used (Taketani et al., 2016). The particle density was assumed to be 2000 kg m^{-3}, and 1 million computational particles were randomly generated in the Arctic region for the backward runs.

Four ground-based observations made during the period 2007–2011 were used to validate the model performance. The potential BC emission sensitivity at 0–500 m above ground level from a 0.1° grid centered at each site was simulated. Other model parameterizations were consistent with those for the Arctic region, except that 200 000 computational particles were released.

2.2 Emission inventories

We focused on BC sources from anthropogenic emissions and open biomass burning. The Hemispheric Transport of Air Pollution version 2 inventory (HTAP2) for 2010 was used for monthly anthropogenic BC emissions (Janssens-Maenhout et al., 2015), which include sectors from energy, industry, residential and transport. It is worth noting that the residential sector includes not only combustions of fossil fuels, but also biofuels. However, as it has been reported that BC emissions in Russia were underestimated in HTAP2, we used the BC emissions reported by Huang et al. (2015) for Russia, in which the annual BC emissions were 224 Gg yr^{-1}. For open biomass burning, we used the monthly BC emissions from the Global Fire Emissions Database version 3 inventory (GFED3) (van der Werf et al., 2010) for the purposes of intercomparison with other studies, as this version is widely used. The term "open biomass burning" here indicates burning of biomass in the open field as is determined by the remote sensing measurement basis, including forest, agricultural waste, peat fires, grassland and savanna, woodland, deforestation and degradation, where biofuel burning for residential use is not included. Geospatial distributions of emissions from anthropogenic sources and open biomass burning in January and July are shown in Fig. S1.

2.3 Calculation of Arctic BC source contributions

The source contributions to Arctic BC were derived by incorporating the gridded retention time into the column emission flux, which was derived from the emission inventories in each grid. Calculations for anthropogenic sources and open biomass burning were performed separately and the sum was used. For anthropogenic sources, the regions were separated into North America and Canada (25–80° N, 50–170° W), Europe (30–80° N, 0–30° E), Russia (53–80° N, 30–180° E), East Asia (35–53° N, 75–150° E and 20–35° N, 100–150° E), and others (the rest) (Fig. 1a). For open biomass burning sources, the regions were

187 separated into Alaska and Canada (50–75° N, 50–170° W), Siberia (50–75° N, 60–180° E),
188 and others (Fig. 1b).
189 2.3 Observations
190 BC levels simulated by Flexpart were compared with those based on surface observations
191 at four sites: Barrow, USA (156.6° W, 71.3° N, 11 m asl), Alert, Canada (62.3° W, 82.5° N,
192 210 m asl), Zeppelin, Norway (11.9° E, 78.9° N, 478 m asl), and Tiksi, Russia (128.9° E, 71.6°
193 N, 8 m asl). Aerosol light absorption was determined by using particle soot absorption
194 photometers (PSAPs) at Barrow, Alert, and Zeppelin, and an aethalometer at Tiksi. For PSAP
195 measurements, the equivalent BC values were derived using a mass absorption efficiency of
196 10 m^2 g^{-1}. The equivalent BC at Tiksi, which was determined with an aethalometer, was
197 obtained directly. These measurement data were obtained from the European Monitoring
198 and Evaluation Programme and World Data Centre for Aerosols database
199 (http://ebas.nilu.no) (Tørseth et al., 2012).
200 It is worth noting that uncertainties could be introduced by using different BC
201 measurement techniques. An evaluation of three methods for measuring BC at Alert,
202 Canada indicated that an average of the refractory BC determined with a single-particle soot
203 photometer (SP2) and elemental carbon (EC) determined from filter samples give the best
204 estimate of BC mass (Sharma et al., 2017). Xu et al. (2017) reported that the equivalent BC
205 determined with a PSAP was close to the average of the values for refractory BC and EC at
206 Alert. In this study, we consider that the equivalent BC values determined with a PSAP at
207 Barrow, Alert, and Zeppelin to be the best estimate. There may be uncertainties in the
208 equivalent BC observations performed with an aethalometer at Tiksi because of co-existing
209 particles such as light-absorptive organic aerosols, scattering particles, and dusts
210 (Kirchstetter et al., 2004; Lack and Langridge, 2013). Interference by the filter and

uncertainties in the mass absorption cross section could also contribute to the bias observed in measurements made with an aethalometer at Tiksi.

3 Results and discussion

3.1. Comparisons of simulations with BC observations at Arctic surface sites

Flexpart generally reproduced the seasonal variations in BC at four Arctic sites well [Pearson correlation coefficient (R) = 0.53–0.80, root-mean-square error (RMSE) = 15.1–56.8 ng m^{-3}] (Fig. 2). Winter maxima were observed for the four sites, while a secondary elevation was observed for Alert and Tiksi. At Barrow, the observed high values of BC were unintentionally excluded during data screening in the forest fire season in summer (Stohl et al., 2013); the original observed BC is supposed to be higher as was reflected by the simulation. This seasonality is probably related to relatively stronger transport to the Arctic region in winter, accompanied by lower BC aging and inefficient removal, as simulated by older versions of Flexpart (Eckhardt et al., 2015; Stohl et al., 2013).

From January to May at Barrow and Alert, the mean BC simulated by Flexpart v10.1 were 32.2 ng m^{-3} and 31.2 ng m^{-3}, respectively. Which was 46 % lower than the observations (59.3 ng m^{-3} and 58.2 ng m^{-3}, respectively). This is probably related to the inadequate BC emission in the inventory, although seasonal variations in residential heating are included in HTAP2, which would reduce the simulation bias (Xu et al., 2017). Simulations by GEOS-Chem using the same emission inventories also underestimated BC levels at Barrow and Alert (Ikeda et al., 2017; Xu et al., 2017). The underestimation by Flexpart could also be partly contributed by the assumption that all particles are hydrophilic, where the BC scavenging could be overestimated. The corresponding uncertainties are larger in winter months, when there are more sources from fossil fuel combustion.

234 At Zeppelin, the Flexpart-simulated BC (39.1 ng m^{-3} for annual mean) was 85 % higher
235 than the observed value (21.1 ng m^{-3} for annual mean), especially in winter (112% higher). It
236 has been reported that riming in mixed-phase clouds occurs frequently at Zeppelin (Qi et al.,
237 2017). During the riming process, BC particles act as ice particles and collide with the
238 relatively numerous water drops, which form frozen cloud droplets, and then snow is
239 precipitated. This results in relatively efficient BC scavenging (Hegg et al., 2011). Such a
240 process could not be dealt with by the model. At Tiksi, Flexpart underestimated BC (74.4 ng
241 m^{-3} for annual mean) in comparison with observation (104.2 ng m^{-3} for annual mean). Other
242 than the hydrophilic BC assumption and underestimated BC emission in the simulation as
243 the cases for Barrow and Alert, the observations at Tiksi by an aethalometer could
244 containing light-absorbing particles other than BC, resulting in higher observed
245 concentrations if compared with those obtained by SP2, EC and PSAP.
246 Anthropogenic emissions are the main sources of BC at the four Arctic sites from late
247 autumn to spring, whereas open biomass burning emissions make large contributions in
248 summer. From October to April, anthropogenic emissions accounted for 87–100 % of BC
249 sources at all the observation sites. At Barrow, open biomass burning accounted for 35–
250 78 % of BC in June–September (Fig. 2). There are large interannual variations in both
251 observed and simulated BC (Fig. S2). In June–August 2010, the mean contributions of open
252 biomass burning to BC were 6.3, 2.4, and 8.6 times those from anthropogenic sources at
253 Alert, Zeppelin, and Tiksi, respectively. In this study, we focused on BC in the Arctic region in
254 2010.
255 **3.2 Potential emission sensitivity of Arctic BC**
256 The potential emission sensitivities (footprint) of Arctic BC showed different patterns
257 with respect to altitude. The Arctic surface is sensitive to emissions at high latitudes (>60°

N). Air masses stayed for over 60 s in each of the 1° grids from the eastern part of northern Eurasia and the Arctic Ocean before being transported to the Arctic surface in the winter, represented by January (Fig. 3a). In comparison, during the summer, represented by July, BC at the Arctic surface was mainly affected by air masses that originated from the Arctic Ocean and the Norwegian Sea (Fig. 3b). These results imply that local BC emissions within the Arctic regions, although relatively weak compared with those from the mid-latitude regions, could strongly affect Arctic air pollution. Local BC emissions are important in the wintertime because the relatively stable boundary layer does not favor pollution dispersion. Recent increases in anthropogenic emissions in the Arctic region, which have been caused by the petroleum industry and development of the Northern Sea Route, are expected to cause deterioration of air quality in the Arctic. Socio-economic developments in the Arctic region would increase local BC emissions, and this will be a non-negligible issue in the future (Roiger et al., 2015; Schmale et al., 2018).

BC at high altitudes in the Arctic is more sensitive to mid-latitude (30–60° N) emissions, especially in wintertime. In January, air masses hovered over the Bering Sea and the North Atlantic Ocean before arriving at the Arctic (Fig. 3c). A notable corridor at 30–50° N covering Eurasia and the United States was the sensitive region that affected BC at high altitudes in the Arctic in January. These results indicate that mid-latitude emissions, especially those with relatively large strengths from East Asia, East America, and Europe, could alter the atmospheric constituents at high altitudes in the Arctic. Central to east Siberia was the most sensitive region for BC at high altitudes in the Arctic in July (Fig. 3d). These results suggest that pollutants from frequent and extensive wildfires in Siberia in summer are readily transported to high altitudes in the Arctic. Boreal fires are expected to occur more frequently and over larger burning areas under future warming (Veira et al., 2016),

Deleted: (~ 5000 m)

283 therefore the atmospheric constituents and climate in the Arctic could undergo more rapid
284 changes.

285 **3.3 Seasonal variations and sources of Arctic surface BC**

286 Arctic surface BC showed clear seasonal variations, with a primary peak in winter–spring
287 (December–March, 61.8–82.8 ng m^{-3}) and a secondary peak in summer (July, 52.7 ng m^{-3}).
288 BC levels were relatively low in May–June (21.8–23.1 ng m^{-3}) and September–November
289 (34.1–40.9 ng m^{-3}) (Fig. 4a). This seasonality agrees with observations and simulations at
290 Alert, Tiksi, and Barrow if consider the unintentional data exclusion (Stohl et al., 2013), and
291 previous studies targeting the whole Arctic (Ikeda et al., 2017; Xu et al., 2017). Compared
292 with the study reported by Stohl et al. (2013), the current work using the new scheme
293 produced smaller discrepancies between the simulated data and observations. Although the
294 simulation periods (monthly means for 2007–2011 in this study and for 2008–2010 in the
295 old scheme) and the anthropogenic emission inventories (HTAP2 in this study and ECLIPSE4
296 in the previous study) are different, the new scheme shows potential for better representing
297 BC transport and removal processes in the Arctic.

298 The annual mean Arctic BC at the surface was estimated to be 48.2 ng m^{-3}. From October
299 to April, anthropogenic sources accounted for 96–100 % of total BC at the Arctic surface.
300 Specifically, anthropogenic emissions from Russia accounted for 61–76 % of total BC in
301 October–May (56 % annually), and was the dominant sources of Arctic BC at the surface.
302 From an isentropic perspective, the meteorological conditions in winter favored the
303 transport of pollutants from northern Eurasia to the lower Arctic, along with diabatic cooling
304 and strong inversions (Klonecki et al., 2003). In comparison, open biomass burning from
305 boreal regions accounted for 56–85 % (75 % on average) of Arctic BC at the surface in
306 summer; open biomass burning emissions from North America and Canada accounted for

307 54 % of total Arctic surface BC in June, and those from Siberia accounted for 59–61 % in
308 July–August. Wildfires in the boreal forests in summer had a major effect on air quality in
309 the Arctic.

310 On an annual basis, anthropogenic sources and open biomass burning emissions
311 accounted for 82 % and 18 %, respectively, of total Arctic surface BC. In which, gas flaring
312 and residential burning (including burning of fossil fuels and biofuels) are accounting for
313 36 % (28–57 % in October–March) and 15 % (13–25 % in October–March), respectively (Fig.
314 5a-b). Our results support Stohl et al. (2013) that residential combustion emissions,
315 especially in winter are important sources of Arctic BC (Table 1). We estimated a
316 contribution of gas flaring to Arctic surface BC of 17.5 ng m^{-3} (36% of total). In comparison,
317 the value was estimated as 11.8 ng m^{-3} using an average Arctic surface BC of 28 ng m^{-3} and
318 a fraction from gas flaring of 42 % evaluated by earlier versions of Flexpart (Stohl et al.,
319 2013; Winiger et al., 2019). The different contribution could be partly attributed to the
320 difference in gas flaring emission inventory. BC emission from gas flaring in Russia by Huang
321 et al. (2015) was used in the current study, where total BC emission from gas flaring in
322 Russia in 2010 was ca. 81.1 kilotonne, which was larger than the estimate of ca. 64.9
323 kilotonne by GAINS inventory (Klimont et al., 2017) used by Stohl et al. (2013). Moreover,
324 Adopting ECLIPSEv5 inventory as was used by Winiger et al. (2019), we estimated that gas
325 flaring was contributing 14.4 ng m^{-3} to Arctic surface BC using Flexpart v10.1, a value 22 %
326 higher than those obtained using Flexpart v9. This difference could be attributed to the
327 improvement of the wet-scavenging scheme by Flexpart v10.1.

328 A recent study based on isotope observations at the Arctic sites and Flexpart v9.2
329 simulation suggested that open biomass burning, including open field burning and
330 residential biofuel burning, contributed 39 % of annual BC in 2011–2015 (Winiger et al.,

2019) (Table 1). In comparison, we estimated that residential burning and open biomass burning together account for 33 % of total Arctic surface BC. As the residential burning in our study includes burning of both biofuels and fossil fuels, our results indicated that biomass burning has a relatively smaller contribution. Other than the differences in BC removal treatment between different versions of the model, the contribution difference could also be attributed to the different emission inventories and years (2010 versus 2011-2015).

The geospatial contributions of anthropogenic sources and open biomass burning emissions can be further illustrated by taking January and July as examples. In January, high levels of anthropogenic emissions from Russia (contributing 64 % of Arctic surface BC), Europe (18 %), and East Asia (9 %) were identified (Fig. 6a). Specifically, Yamalo-Nenets Autonomous Okrug in Russia, which has the largest reserves of Russia's natural gas and oil (Filimonova et al., 2018), was the most notable emission hotspot, which suggests gas-flaring sources. The Komi Republic in Russia was also identified as a strong anthropogenic emitter contributing to Arctic surface BC. These gas-flaring industrial regions in Russia (58–69° N, 68–81°E) together contributed 33 % and 31 % of Arctic surface BC for January and the annual mean, respectively. Recently, Dong et al. (2019) evaluated BC emission inventories using GEOS-Chem and proposed that using the inventory compiled by Huang et al. (2015) for Russia, in which gas flaring accounted for 36 % of anthropogenic emissions, had no prominent impact on the simulation performance in Russia and the Arctic. They suggested that use of a new global inventory for BC emissions from natural gas flaring would improve the model performance (Huang and Fu, 2016). These results suggest that inclusion of BC emissions from gas flaring on the global scale is necessary for further BC simulations.

In Europe, a relatively high contribution of anthropogenic emissions to Arctic surface BC in January was made by Poland (50–55° N, 15–24° E, contributing 4 % of Arctic surface BC) because of relatively large emission fluxes in the region (Fig. S1a). Anthropogenic emissions from East China, especially those north of ~33° N (33–43° N, 109–126° E), contributed perceptibly (5 %) to Arctic surface BC.

In July, contributions from anthropogenic sources shrank to those from Yamalo-Nenets Autonomous Okrug and Komi Republic in Russia, and contributed a lower fraction (3 % of Arctic surface BC) (Fig. 6b). Few open biomass burning sources contributed in January (Fig. 6c), but contributions from open biomass burning to Arctic surface BC in July can be clearly seen, mainly from the far east of Russia, Canada, and Alaska (Fig. 6d). Open biomass burning emissions from Kazakhstan, southwest Russia, southern Siberia, and northeast China also contributed to Arctic surface BC, although at relatively low strengths (Fig. 5d and Fig. S1d). However, the contributions from open biomass burning could be higher, as the MODIS burned area, the basis of GFED emission inventories, was underestimated for northern Eurasia by 16 % (Zhu et al., 2017). Evangeliou et al. (2016) estimated a relatively high transport efficiency of BC from open biomass burning emissions to the Arctic, which led to a high contribution, i.e., 60 %, from such sources to BC deposition in the Arctic in 2010. A recent study suggested that open fires burned in western Greenland in summer (31 July to 21 August 2017) could potentially alter the Arctic air composition and foster glacier melting (Evangeliou et al., 2019). Although the footprint of Arctic surface BC showed a relatively weak sensitivity to areas such as forests and tundra, in the boreal regions, pollutants from boreal wildfires could have greater effects on the Arctic air composition in summer under future warming scenarios (Veira et al., 2016).

3.4 Sources of Arctic BC at high altitudes

Arctic BC levels at high altitudes showed the highest levels in spring (March–April, 40.5–53.9 ng m^{-3}), followed by those in late autumn to early winter (November–January, 36.5–40.0 ng m^{-3}), and summer (July–August, 33.0–39.0 ng m^{-3}) (Fig. 4c). The annual mean Arctic BC at high altitudes was estimated to be 35.2 ng m^{-3}, which is ca. 73 % of those at the surface. Such a vertical profile is in accordance with those based on aircraft measurements over the High Canadian Arctic (Schulz et al., 2019). Similarly to the case for the surface, anthropogenic sources dominated by residential sectors, transport, industry and energy (excluding gas flaring), accounted for 94–100 % of Arctic BC at high altitudes in October–May (Figs. 4c, 5c). East Asia accounted for 34–65 % of the total BC in October–May (40 % annually). In comparison, using the Community Atmosphere Model version 5 driven by the NASA Modern Era Retrospective-Analysis for Research and Applications reanalysis data and the IPCC AR5 year 2000 BC emission inventory, H. Wang et al. (2014) found that East Asia accounted for 23% of BC burden in the Arctic for 1995–2005. In summer, open biomass burning in the boreal regions accounted for 40–72 % (57 % on average) of Arctic BC at high altitudes, similar to the source contributions to Arctic surface BC. Specifically, open biomass burning sources from Siberia accounted for 40–42 % of Arctic BC at high altitudes in July–August. Annually, anthropogenic sources and open biomass burning accounted for 83 % (in which residential sources accounted for 34%) and 17 %, respectively, of total Arctic BC at high altitudes (Figs. 4d, 5d).

Further investigations of geospatial contributions to Arctic BC at high altitudes in January and July provided more details regarding BC sources. In January, the main anthropogenic BC source in East Asia covered a wide range in China (Fig. 7a). Not only east and northeast China, but also southwest China (Sichuan and Guizhou provinces) were the major anthropogenic sources of Arctic BC at high altitudes. In July, anthropogenic sources made a

Deleted: (4750–4250 m)

relatively weak contribution to Arctic BC at high altitudes. The regions that were sources of open biomass burning contributions to Arctic BC at high altitudes were mainly the far east of Siberia, Kazakhstan, central Canada, and Alaska, i.e., similar to the sources of Arctic surface BC. Unlike Arctic surface BC, for which the dominant source regions are at high latitudes in both winter and summer, Arctic BC at high altitudes mainly originates from mid-latitude regions (Figs. 6 and 7). In terms of transport pathways, air masses could be uplifted at low-to-mid latitudes and transported to the Arctic (Stohl, 2006). Further investigations are needed to obtain more details of the transport processes.

3.5 Comparison of Flexpart and GEOS-Chem simulations of BC sources

Data for BC sources simulated with Flexpart were compared with those obtained with GEOS-Chem (Ikeda et al., 2017), which is an Eulerian atmospheric transport model, using the same emission inventories. The simulated seasonal variations in Arctic BC levels and source contributions obtained with Flexpart agreed well with those obtained with GEOS-Chem (Fig. S3). The annual mean BC levels at the Arctic surface obtained by Flexpart and GEOS-Chem simulations were 48 and 70 ng m^{-3}, respectively; the high-altitude values simulated by Flexpart and GEOS-Chem were 35 and 38 ng m^{-3}, respectively. The magnitude difference between the BC levels at the Arctic surface could be related to meteorology. ECMWF ERA-Interim data were used as the input for the Flexpart simulation, whereas the GEOS-Chem simulation was driven by assimilated meteorological data from the Goddard Earth Observation System (GEOS-5).

The treatments of the BC removal processes could also lead to different simulation results, depending on the model. In terms of BC loss processes, dry and wet depositions were the removal pathways, depending on the particle size and density, in Flexpart. The treatment of meteorology, especially cloud water and precipitation, would therefore affect

the uncertainties of the simulations. In Flexpart version 10.1, BC particles are separately parameterized as ice nuclei for ice clouds, cloud condensation nuclei for liquid-water clouds, and 90 % as cloud condensation nuclei for mixed-phase clouds. The separation of mixed-phase clouds is realistic, as 77 % of in-cloud scavenging processes occurred in the mixed phase over a 90 day period starting from December 2006 (Grythe et al., 2017).

In GEOS-Chem simulations, the BC aging was parameterized based on the number concentration of OH radicals (Liu et al., 2011). The BC was assumed to be hydrophilic in liquid clouds ($T \geq 258$ K) and hydrophobic when serving as ice nuclei in ice clouds ($T < 258$ K) (Wang et al., 2011), with modifications because the scavenging rate of hydrophobic BC was reduced to 5 % of water-soluble aerosols for liquid clouds (Bourgeois and Bey, 2011). Such a treatment is expected to improve the simulation accuracy (Ikeda et al., 2017).

In Lagrangian models, the trajectories of particles are computed by following the movement of air masses with no numerical diffusion, although some artificial numerical errors could be generated from stochastic differential equations (Ramli and Esler, 2016). As a result, long-range transport processes can be well simulated (Stohl, 2006; Stohl et al., 2013). In comparison, Eulerian chemical transport models such as GEOS-Chem have the advantage of simulating non-linear processes on the global scale, which enables treatment of the BC aging processes (coating with soluble components) (Bey et al, 2001; Eastham et al., 2018). However, with GEOS-Chem, the capture of intercontinental pollution plumes is difficult because of numerical plume dissipation (Rastigejev et al., 2010). Nevertheless, the agreement between the Flexpart and GEO-Chem simulations of BC source contributions indicates improved reliability of evaluated source contributions to Arctic BC.

4 Conclusions

The source contributions to Arctic BC were investigated by using a Flexpart (version 10.1) transport model that incorporated emission inventories. Flexpart-simulated BC data agreed well with observations at Arctic sites, i.e., Barrow, Alert, Zeppelin, and Tiksi. The source regions and source sectors of BC at the surface and high altitudes over a wide region in the Arctic north of 66° N were simulated. BC at the Arctic surface was sensitive to local emissions and those from nearby Nordic countries (>60° N). These results emphasize the role of anthropogenic emissions such as gas flaring and development of the Northern Sea Route in affecting air quality and climate change in the Arctic. Anthropogenic emissions in the northern regions of Russia were the main source (56 %) of Arctic surface BC annually. In contrast, BC in the Arctic at high altitudes was sensitive to mid-latitude emissions (30–60° N). Although they are geospatially far from the Arctic, anthropogenic emissions in East Asia made a notable (40 %) contribution to BC in the Arctic at high altitudes annually. Open biomass burning emissions, which were mainly from Siberia, Alaska, and Canada, were important in summer, contributing 56–85 % of BC at the Arctic surface, and 40–72 % at Arctic high altitudes. Future increases in wildfires as a result of global warming could therefore increase the air pollution level during the Arctic summer. This study clarifies the source regions and sectors of BC in the Arctic. This information is fundamental for understanding and tackling air pollution and climate change in the region.

Deleted: (~ 5000 m)

Data Availability. The data set for simulated footprint and BC source contributions is available on request to the corresponding author.

Author contributions. CZ and YK designed the study. CZ, MT, and IP optimized the Flexpart model. CZ performed Flexpart model simulations, conducted analyses, and wrote the

manuscript. KI and HT provided data for GEOS-Chem simulations and site observations. All authors made comments that improved the paper.

Competing interests. The authors declare that they have no conflict of interest.

Financial Support. This study was supported by the Environmental Research and Technology Development Fund (2-1505) of the Ministry of the Environment, Japan.

Acknowledgment. We acknowledge staffs from the following university and agencies for BC observational data: Barrow and Tiksi sites are operated by National Oceanic and Atmospheric Administration; Zeppelin site is operated by Stockholm University; and Alert site is operated by Environment and Climate Change Canada. We are grateful to Chandra Mouli Pavuluri (Tianjin University) and an anonymous reviewer for the comments. We thank Helen McPherson, PhD, from Edanz Group (www.edanzediting.com/ac) for editing a draft of this manuscript.

Deleted: two

Deleted: s

References

AMAP Assessment 2015: Black carbon and ozone as Arctic climate forcers, Arctic Monitoring and Assessment Programme (AMAP), Oslo, Norway, 2015.

Bey, I., Jacob, D. J., Yantosca, R. M., Logan, J. A., Field, B. D., Fiore, A. M., Li, Q. B., Liu, H. G. Y., Mickley, L. J., and Schultz, M. G.: Global modeling of tropospheric chemistry with assimilated meteorology: Model description and evaluation, J Geophys Res-Atmos, 106, 23073-23095, doi:10.1029/2001jd000807, 2001.

Bourgeois, Q. and Bey, I.: Pollution transport efficiency toward the Arctic: sensitivity to aerosol scavenging and source regions, J. Geophys. Res., 116, D08213, doi:10.1029/2010JD015096, 2011.

503 Brock, C. A., Cozic, J., Bahreini, R., Froyd, K. D., Middlebrook, A. M., McComiskey, A.,
504 Brioude, J., Cooper, O. R., Stohl, A., Aikin, K. C., de Gouw, J. A., Fahey, D. W., Ferrare, R.
505 A., Gao, R. S., Gore, W., Holloway, J. S., Hubler, G., Jefferson, A., Lack, D. A., Lance, S.,
506 Moore, R. H., Murphy, D. M., Nenes, A., Novelli, P. C., Nowak, J. B., Ogren, J. A., Peischl, J.,
507 Pierce, R. B., Pilewskie, P., Quinn, P. K., Ryerson, T. B., Schmidt, K. S., Schwarz, J. P.,
508 Sodemann, H., Spackman, J. R., Stark, H., Thomson, D. S., Thornberry, T., Veres, P., Watts,
509 L. A., Warneke, C., and Wollny, A. G.: Characteristics, sources, and transport of aerosols
510 measured in spring 2008 during the aerosol, radiation, and cloud processes affecting
511 Arctic Climate (ARCPAC) Project, Atmospheric Chemistry and Physics, 11, 2423-2453,
512 doi:10.5194/acp-11-2423-2011, 2011.
513 Cohen, J., Screen, J. A., Furtado, J. C., Barlow, M., Whittleston, D., Coumou, D., Francis, J.,
514 Dethloff, K., Entekhabi, D., Overland, J., and Jones, J.: Recent Arctic amplification and
515 extreme mid-latitude weather, Nature Geoscience, 7, 627-637, doi:10.1038/Ngeo2234,
516 2014.
517 Cozic, J., Verheggen, B., Mertes, S., Connolly, P., Bower, K., Petzold, A., Baltensperger, U.,
518 and Weingartner, E.: Scavenging of black carbon in mixed phase clouds at the high alpine
519 site Jungfraujoch, Atmos. Chem. Phys., 7, 1797-1807, doi:10.5194/acp-7-1797-2007,
520 2007.
521 Dong, X., Zhu, Q., Fu, J. S., Huang, K., Tan, J., and Tipton, M.: Evaluating recent updated black
522 carbon emissions and revisiting the direct radiative forcing in Arctic, Geophysical
523 Research Letters, 46, 3560– 3570. doi:10.1029/2018GL081242, 2019.
524 Eastham, S. D., Long, M. S., Keller, C. A., Lundgren, E., Yantosca, R. M., Zhuang, J. W., Li, C.,
525 Lee, C. J., Yannetti, M., Auer, B. M., Clune, T. L., Kouatchou, J., Putman, W. M., Thompson,
526 M. A., Trayanov, A. L., Molod, A. M., Martin, R. V., and Jacob, D. J.: GEOS-Chem High
527 Performance (GCHP v11-02c): a next-generation implementation of the GEOS-Chem
528 chemical transport model for massively parallel applications, Geosci Model Dev, 11,
529 2941-2953, doi:10.5194/gmd-11-2941-2018, 2018.
530 Eckhardt, S., Quennehen, B., Olivie, D. J. L., Berntsen, T. K., Cherian, R., Christensen, J. H.,
531 Collins, W., Crepinsek, S., Daskalakis, N., Flanner, M., Herber, A., Heyes, C., Hodnebrog,
532 O., Huang, L., Kanakidou, M., Klimont, Z., Langner, J., Law, K. S., Lund, M. T., Mahmood,
533 R., Massling, A., Myriokefalitakis, S., Nielsen, I. E., Nojgaard, J. K., Quaas, J., Quinn, P. K.,
534 Raut, J. C., Rumbold, S. T., Schulz, M., Sharma, S., Skeie, R. B., Skov, H., Uttal, T., von

535 Salzen, K., and Stohl, A.: Current model capabilities for simulating black carbon and
536 sulfate concentrations in the Arctic atmosphere: a multi-model evaluation using a
537 comprehensive measurement data set, Atmospheric Chemistry and Physics, 15, 9413-
538 9433, doi:10.5194/acp-15-9413-2015, 2015.
539 Evangeliou, N., Balkanski, Y., Hao, W. M., Petkov, A., Silverstein, R. P., Corley, R., Nordgren,
540 B. L., Urbanski, S. P., Eckhardt, S., Stohl, A., Tunved, P., Crepinsek, S., Jefferson, A.,
541 Sharma, S., Nojgaard, J. K., and Skov, H.: Wildfires in northern Eurasia affect the budget of
542 black carbon in the Arctic - a 12-year retrospective synopsis (2002-2013), Atmospheric
543 Chemistry and Physics, 16, 7587-7604, doi:10.5194/acp-16-7587-2016, 2016.
544 Evangeliou, N., Kylling, A., Eckhardt, S., Myroniuk, V., Stebel, K., Paugam, R., Zibtsev, S., and
545 Stohl, A.: Open fires in Greenland in summer 2017: transport, deposition and radiative
546 effects of BC, OC and BrC emissions, Atmospheric Chemistry and Physics, 19, 1393-1411,
547 doi:10.5194/acp-19-1393-2019, 2019.
548 Filimonova, I. V., Komarova, A. V., Eder, L. V., and Provornaya, I. V.: State instruments for the
549 development stimulation of Arctic resources regions, IOP Conference Series: Earth and
550 Environmental Science, 193, 012069, doi:10.1088/1755-1315/193/1/012069, 2018.
551 Garrett, T. J., Brattstrom, S., Sharma, S., Worthy, D. E. J., and Novelli, P.: The role of
552 scavenging in the seasonal transport of black carbon and sulfate to the Arctic,
553 Geophysical Research Letters, 38, L16805, doi:10.1029/2011gl048221, 2011.
554 Grythe, H., Kristiansen, N. I., Zwaaftink, C. D. G., Eckhardt, S., Strom, J., Tunved, P., Krejci, R.,
555 and Stohl, A.: A new aerosol wet removal scheme for the Lagrangian particle model
556 FLEXPART v10, Geosci Model Dev, 10, 1447-1466, doi:10.5194/gmd-10-1447-2017, 2017.
557 Hegg, D. A., Clarke, A. D., Doherty, S. J., and Ström, J.: Measurements of black carbon
558 aerosol washout ratio on Svalbard, Tellus B, 63, 891–900, doi:10.1111/j.1600-
559 0889.2011.00577.x, 2011.
560 Hertel, O., Christensen, J. Runge, E. H., Asman, W. A. H., Berkowicz, R., Hovmand, M. F., and
561 Hov, O.: Development and testing of a new variable scale air pollution model – ACDEP,
562 Atmos. Environ., 29, 1267–1290, 1995.
563 Hirdman, D., Burkhart, J. F., Sodemann, H., Eckhardt, S., Jefferson, A., Quinn, P. K., Sharma,
564 S., Strom, J., and Stohl, A.: Long-term trends of black carbon and sulphate aerosol in the
565 Arctic: changes in atmospheric transport and source region emissions, Atmospheric
566 Chemistry and Physics, 10, 9351-9368, doi:10.5194/acp-10-9351-2010, 2010.

Huang, K., and Fu, J. S.: A global gas flaring black carbon emission rate dataset from 1994 to 2012, Scientific Data, 3, 160104. doi:10.1038/sdata.2016.104, 2016.

Huang, K., Fu, J. S., Prikhodko, V. Y., Storey, J. M., Romanov, A., Hodson, E. L., Cresko, J., Morozova, I., Ignatieva, Y., and Cabaniss, J.: Russian anthropogenic black carbon: Emission reconstruction and Arctic black carbon simulation, J Geophys Res-Atmos, 120, 11306-11333, doi:10.1002/2015jd023358, 2015.

Ikeda, K., Tanimoto, H., Sugita, T., Akiyoshi, H., Kanaya, Y., Zhu, C. M., and Taketani, F.: Tagged tracer simulations of black carbon in the Arctic: transport, source contributions, and budget, Atmospheric Chemistry and Physics, 17, 10515-10533, doi:10.5194/acp-17-10515-2017, 2017.

Janssens-Maenhout, G., Crippa, M., Guizzardi, D., Dentener, F., Muntean, M., Pouliot, G., Keating, T., Zhang, Q., Kurokawa, J., Wankmüller, R., Denier van der Gon, H., Kuenen, J. J. P., Klimont, Z., Frost, G., Darras, S., Koffi, B., and Li, M.: HTAP_v2.2: a mosaic of regional and global emission grid maps for 2008 and 2010 to study hemispheric transport of air pollution, Atmos. Chem. Phys., 15, 11411–11432, doi:10.5194/acp-15-11411-2015, 2015.

Keegan, K. M., Albert, M. R., McConnell, J. R., and Baker, I.: Climate change and forest fires synergistically drive widespread melt events of the Greenland Ice Sheet, P Natl Acad Sci USA, 111, 7964-7967, doi:10.1073/pnas.1405397111, 2014.

Kipling, Z., Stier, P., Schwarz, J. P., Perring, A. E., Spackman, J. R., Mann, G. W., Johnson, C. E., and Telford, P. J.: Constraints on aerosol processes in climate models from vertically-resolved aircraft observations of black carbon, Atmos. Chem. Phys., 13, 5969–5986, doi:10.5194/acp-13-5969-2013, 2013.

Kirchstetter, T. W., Novakov, T., and Hobbs, P. V.: Evidence that the spectral dependence of light absorption by aerosols is affected by organic carbon, J Geophys Res-Atmos, 109, D21208, doi:10.1029/2004jd004999, 2004.

Klimont, Z., Kupiainen, K., Heyes, C., Purohit, P., Cofala, J., Rafaj, P., Borken-Kleefeld, J., and Schöpp, W.: Global anthropogenic emissions of particulate matter including black carbon, Atmos. Chem. Phys., 17, 8681–8723, doi:10.5194/acp-17- 8681-2017, 2017.

Klonecki, A., Hess, P., Emmons, L., Smith, L., Orlando, J., and Blake, D.: Seasonal changes in the transport of pollutants into the Arctic troposphere-model study, J Geophys Res-Atmos, 108, 8367, doi:10.1029/2002jd002199, 2003.

598 Koch, D., and Hansen, J.: Distant origins of Arctic black carbon: A Goddard Institute for Space
599 Studies ModelE experiment, J Geophys Res-Atmos, 110, D04204,
600 doi:10.1029/2004jd005296, 2005.
601 Koch, D., Schmidt, G. A., and Field, C. V.: Sulfur, sea salt, and radionuclide aerosols in GISS
602 ModelE, J. Geophys. Res., 111, D06206, doi:10.1029/2004jd005550, 2006.
603 Lack, D. A., and Langridge, J. M.: On the attribution of black and brown carbon light
604 absorption using the Angstrom exponent, Atmospheric Chemistry and Physics, 13,
605 10535–10543, doi:10.5194/acp-13-10535-2013, 2013.
606 Law, K. S., and Stohl, A.: Arctic air pollution: Origins and impacts, Science, 315, 1537–1540,
607 doi:10.1126/science.1137695, 2007.
608 Liu, J., Fan, S., Horowitz, L. W., and Levy II, H.: Evaluation of factors controlling long-range
609 transport of black carbon to the Arctic, J. Geophys. Res., 116, D00A14,
610 doi:10.1029/2010JD015145, 2011.
611 Liu, D., Quennehen, B., Darbyshire, E., Allan, J. D., Williams, P. I., Taylor, J. W., Bauguitte, S. J.
612 B., Flynn, M. J., Lowe, D., Gallagher, M. W., Bower, K. N., Choularton, T. W., and Coe, H.:
613 The importance of Asia as a source of black carbon to the European Arctic during
614 springtime 2013, Atmospheric Chemistry and Physics, 15, 11537–11555, doi:10.5194/acp-
615 15-11537-2015, 2015.
616 Ma, P.-L., Rasch, P. J., Wang, H., Zhang, K., Easter, R. C., Tilmes, S., Fast, J. D., Liu, X., Yoon, J.-
617 H., and Lamarque, J.-F.: The role of circulation features on black carbon transport into the
618 Arctic in the Community Atmosphere Model version 5 (CAM5), J. Geophys. Res.-Atmos.,
619 118, 4657–4669, doi:10.1002/jgrd.50411, 2013.
620 McMahon, T. A. and Denison, P. J.: Empirical atmospheric deposition parameters – a survey,
621 Atmos. Environ., 13, 571–585, doi:10.1016/0004-6981(79)90186-0, 1979.
622 Najafi, M. R., Zwiers, F. W., and Gillett, N. P.: Attribution of Arctic temperature change to
623 greenhouse-gas and aerosol influences, Nature Climate Change, 5, 246-249,
624 doi:10.1038/Nclimate2524, 2015.
625 Park, R. J., Jacob, D. J., Palmer, P. I., Clarke, A. D., Weber, R. J., Zondlo, M. A., Eisele, F. L.,
626 Bandy, A. R., Thornton, D. C., Sachse, G. W., and Bond, T. C.: Export efficiency of black
627 carbon aerosol in continental outflow: Global implications, J Geophys Res-Atmos, 110,
628 D11205, doi:10.1029/2004jd005432, 2005.

629 Qi, L., Li, Q. B., Henze, D. K., Tseng, H. L., and He, C. L.: Sources of springtime surface black
630 carbon in the Arctic: an adjoint analysis for April 2008, Atmospheric Chemistry and
631 Physics, 17, 9697–9716, doi:10.5194/acp-17-9697-2017, 2017.
632 Quinn, P. K., Bates, T. S., Baum, E., Doubleday, N., Fiore, A. M., Flanner, M., Fridlind, A.,
633 Garrett, T. J., Koch, D., Menon, S., Shindell, D., Stohl, A., and Warren, S. G.: Short-lived
634 pollutants in the Arctic: their climate impact and possible mitigation strategies,
635 Atmospheric Chemistry and Physics, 8, 1723–1735, doi:10.5194/acp-8-1723-2008, 2008.
636 Ramli, H. M. and Esler, J. G.: Quantitative evaluation of numerical integration schemes for
637 Lagrangian particle dispersion models, Geosci. Model Dev., 9, 2441–2457,
638 doi:10.5194/gmd-9-2441-2016, 2016.
639 Rastigejev, Y., Park, R., Brenner, M., and Jacob, D.: Resolving intercontinental pollution
640 plumes in global models of atmospheric transport, J. Geophys. Res., 115, D02302,
641 doi:10.1029/2009JD012568, 2010.
642 Roiger, A., Thomas, J. L., Schlager, H., Law, K. S., Kim, J., Schafler, A., Weinzierl, B.,
643 Dahloktter, F., Krisch, I., Marelle, L., Minikin, A., Raut, J. C., Reiter, A., Rose, M., Scheibe,
644 M., Stock, P., Baumann, R., Bouapar, I., Clerbaux, C., George, M., Onishi, I., and Flemming,
645 J.: Quantifying emerging local anthropogenic emissions in the Arctic region: The ACCESS
646 aircraft campaign experiment, B. Am. Meteorol. Soc., 96, 441–460, doi:10.1175/Bams-D-
647 13-00169.1, 2015.
648 Sand, M., Berntsen, T. K., von Salzen, K., Flanner, M. G., Langner, J., and Victor, D. G.:
649 Response of Arctic temperature to changes in emissions of short-lived climate forcers,
650 Nature Climate Change, 6, 286-289, doi:10.1038/Nclimate2880, 2016.
651 Schacht, J., Heinold, B., Quaas, J., Backman, J., Cherian, R., Ehrlich, A., Herber, A., Huang, W.
652 T. K., Kondo, Y., Massling, A., Sinha, P. R., Weinzierl, B., Zanatta, M., and Tegen, I.: The
653 importance of the representation of air pollution emissions for the modeled distribution
654 and radiative effects of black carbon in the Arctic, Atmos. Chem. Phys., 19, 11159–11183,
655 doi:10.5194/acp-19-11159-2019, 2019.
656 Schmale, J., Arnold, S. R., Law, K. S., Thorp, T., Anenberg, S., Simpson, W. R., Mao, J., and
657 Pratt, K. A.: Local Arctic Air Pollution: A Neglected but Serious Problem, Earths Future, 6,
658 1385–1412, doi:10.1029/2018ef000952, 2018.
659 Schulz, H., Zanatta, M., Bozem, H., Leaitch, W. R., Herber, A. B., Burkart, J., Willis, M. D.,
660 Kunkel, D., Hoor, P. M., Abbatt, J. P. D., and Gerdes, R.: High Arctic aircraft measurements

characterising black carbon vertical variability in spring and summer, Atmospheric Chemistry and Physics, 19, 2361–2384, doi:10.5194/acp-19-2361-2019, 2019.

Sharma, S., Leaitch, W. R., Huang, L., Veber, D., Kolonjari, F., Zhang, W., Hanna, S. J., Bertram, A. K., and Ogren, J. A.: An evaluation of three methods for measuring black carbon in Alert, Canada, Atmospheric Chemistry and Physics, 17, 15225–15243, doi:10.5194/acp-17-15225-2017, 2017.

Shaw, G. E.: The arctic haze phenomenon, B Am Meteorol Soc, 76, 2403–2413, 1995.

Shen, Z. Y., Ming, Y., Horowitz, L. W., Ramaswamy, V., and Lin, M. Y.: On the seasonality of Arctic black carbon, J. Climate, 30, 4429–4441, doi:10.1175/Jcli-D-16-0580.1, 2017.

Shindell, D. T., Chin, M., Dentener, F., Doherty, R. M., Faluvegi, G., Fiore, A. M., Hess, P., Koch, D. M., MacKenzie, I. A., Sanderson, M. G., Schultz, M. G., Schulz, M., Stevenson, D. S., Teich, H., Textor, C., Wild, O., Bergmann, D. J., Bey, I., Bian, H., Cuvelier, C., Duncan, B. N., Folberth, G., Horowitz, L. W., Jonson, J., Kaminski, J. W., Marmer, E., Park, R., Pringle, K. J., Schroeder, S., Szopa, S., Takemura, T., Zeng, G., Keating, T. J., and Zuber, A.: A multi-model assessment of pollution transport to the Arctic, Atmos. Chem. Phys., 8, 5353-5372, doi:10.5194/acp-8-5353-2008, 2008.

Sinha, P. R., Kondo, Y., Koike, M., Ogren, J. A., Jefferson, A., Barrett, T. E., Sheesley, R. J., Ohata, S., Moteki, N., Coe, H., Liu, D., Irwin, M., Tunved, P., Quinn, P. K., and Zhao, Y.: Evaluation of ground-based black carbon measurements by filter-based photometers at two Arctic sites, J Geophys Res-Atmos, 122, 3544-3572, doi:10.1002/2016jd025843, 2017.

Stohl, A., Forster, C., Frank, A., Seibert, P., and Wotawa, G.: Technical note: The Lagrangian particle dispersion model FLEXPART version 6.2, Atmos. Chem. Phys., 5, 2461–2474, doi:10.5194/acp-5-2461-2005, 2005.

Stohl, A.: Characteristics of atmospheric transport into the Arctic troposphere, J Geophys Res-Atmos, 111, D11306, doi:10.1029/2005jd006888, 2006.

Stohl, A., Hittenberger, M., and Wotawa, G.: Validation of the Lagrangian particle dispersion model FLEXPART against large-scale tracer experiment data, Atmos Environ, 32, 4245-4264, doi:10.1016/S1352-2310(98)00184-8, 1998.

Stohl, A., Klimont, Z., Eckhardt, S., Kupiainen, K., Shevchenko, V. P., Kopeikin, V. M., and Novigatsky, A. N.: Black carbon in the Arctic: the underestimated role of gas flaring and

residential combustion emissions, Atmospheric Chemistry and Physics, 13, 8833-8855, doi:10.5194/acp-13-8833-2013, 2013.

Taketani, F., Miyakawa, T., Takashima, H., Komazaki, Y., Pan, X., Kanaya, Y., and Inoue, J.: Shipborne observations of atmospheric black carbon aerosol particles over the Arctic Ocean, Bering Sea, and North Pacific Ocean during September 2014, J Geophys Res-Atmos, 121, 1914-1921, doi:10.1002/2015jd023648, 2016.

Tietze, K., Riedi, J., Stohl, A., and Garrett, T. J.: Space-based evaluation of interactions between aerosols and low-level Arctic clouds during the Spring and Summer of 2008, Atmospheric Chemistry and Physics, 11, 3359-3373, doi:10.5194/acp-11-3359-2011, 2011.

Tilling, R. L., Ridout, A., Shepherd, A., and Wingham, D. J.: Increased Arctic sea ice volume after anomalously low melting in 2013, Nature Geoscience, 8, 643-646, doi:10.1038/Ngeo2489, 2015.

Tørseth, K., Aas, W., Breivik, K., Fjæraa, A. M., Fiebig, M., Hjellbrekke, A. G., Lund Myhre, C., Solberg, S., and Yttri, K. E.: Introduction to the European Monitoring and Evaluation Programme (EMEP) and observed atmospheric composition change during 1972–2009, Atmos. Chem. Phys., 12, 5447–5481, doi:10.5194/acp-12-5447-2012, 2012.

Trusel, L. D., Das, S. B., Osman, M. B., Evans, M. J., Smith, B., Fettweis, X., McConnell, J. R., Noel, B. P. Y., and van den Broeke, M. R.: Nonlinear rise in Greenland runoff in response to post-industrial Arctic warming, Nature, 564, 104-108, doi:10.1038/s41586-018-0752-4, 2018.

van der Werf, G. R., Randerson, J. T., Giglio, L., Collatz, G. J., Mu, M., Kasibhatla, P. S., Morton, D. C., DeFries, R. S., Jin, Y., and van Leeuwen, T. T.: Global fire emissions and the contribution of deforestation, savanna, forest, agricultural, and peat fires (1997–2009), Atmospheric Chemistry and Physics, 10, 11707-11735, doi:10.5194/acp-10-11707-2010, 2010.

Veira, A., Lasslop, G., and Kloster, S.: Wildfires in a warmer climate: Emission fluxes, emission heights, and black carbon concentrations in 2090-2099, J Geophys Res-Atmos, 121, 3195-3223, doi:10.1002/2015jd024142, 2016.

Wang, H., Rasch, P. J., Easter, R. C., Singh, B., Zhang, R., Ma, P. L., Qian, Y., and Beagley, N.: Using an explicit emission tag- ging method in global modeling of source-receptor

relationships for black carbon in the Arctic: Variations, Sources and Trans- port pathways, J. Geophys. Res.-Atmos., 119, 12888–12909, doi:10.1002/2014JD022297, 2014.

Wang, Q., Jacob, D. J., Fisher, J. A., Mao, J., Leibensperger, E. M., Carouge, C. C., Le Sager, P., Kondo, Y., Jimenez, J. L., Cubison, M. J., and Doherty, S. J.: Sources of carbonaceous aerosols and deposited black carbon in the Arctic in winter-spring: implications for radiative forcing, Atmos. Chem. Phys., 11, 12453– 12473, doi:10.5194/acp-11-12453-2011, 2011.

Wang, Q., Jacob, D. J., Spackman, J. R., Perring, A. E., Schwarz, J. P., Moteki, N., Marais, E. A., Ge, C., Wang, J., and Barrett, S. R. H.: Global budget and radiative forcing of black carbon aerosol: constraints from pole-to-pole (HIPPO) observations across the Pacific, J. Geophys. Res. Atmos., 119, 195–206, doi:10.1002/2013JD020824, 2014.

Winiger, P., Andersson, A., Eckhardt, S., Stohl, A., and Gustafsson, O.: The sources of atmospheric black carbon at a European gateway to the Arctic, Nat Commun, 7, 12776, doi:10.1038/ncomms12776, 2016.

Winiger, P., Barrett, T. E., Sheesley, R. J., Huang, L., Sharma, S., Barrie, L. A., Yttri, K. E., Evangeliou, N., Eckhardt, S., Stohl, A., Klimont, Z., Heyes, C., Semiletov, I. P., Dudarev, O. V., Charkin, A., Shakhova, N., Holmstrand, H., Andersson, A., and Gustafsson, O.: Source apportionment of circum-Arctic atmospheric black carbon from isotopes and modeling, Sci Adv, 5, eaau8052, doi:10.1126/sciadv.aau8052, 2019.

Xu, J. W., Martin, R. V., Morrow, A., Sharma, S., Huang, L., Leaitch, W. R., Burkart, J., Schulz, H., Zanatta, M., Willis, M. D., Henze, D. K., Lee, C. J., Herber, A. B., and Abbatt, J. P. D.: Source attribution of Arctic black carbon constrained by aircraft and surface measurements, Atmospheric Chemistry and Physics, 17, 11971-11989, doi:10.5194/acp-17-11971-2017, 2017.

Yang, Y., Wang, H., Smith, S. J., Easter, R. C., and Rasch, P. J.: Sulfate aerosol in the Arctic: Source attribution and radiative forcing, J. Geophys. Res. Atmos., 123, 1899–1918, doi:10.1002/2017JD027298, 2018.

Yang, Y., Smith, S. J., Wang, H., Mills, C. M., and Rasch, P. J.: Variability, timescales, and nonlinearity in climate responses to black carbon emissions, Atmospheric Chemistry and Physics, 19, 2405–2420, doi:10.5194/acp-19-2405-2019, 2019.

Yu, K., Keller, C. A., Jacob, D. J., Molod, A. M., Eastham, S. D., and Long, M. S.: Errors and improvements in the use of archived meteorological data for chemical transport

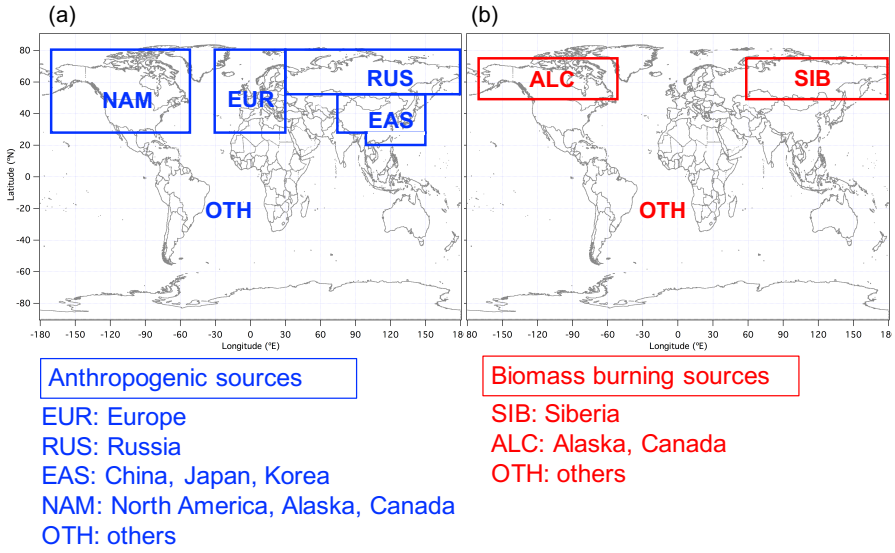
755 modeling: an analysis using GEOS-Chem v11-01 driven by GEOS-5 meteorology, Geosci
756 Model Dev, 11, 305-319, doi:10.5194/gmd-11-305-2018, 2018.
757 Zhu, C., Kobayashi, H., Kanaya, Y., and Saito, M.: Size-dependent validation of MODIS
758 MCD64A1 burned area over six vegetation types in boreal Eurasia: Large underestimation
759 in croplands, Scientific reports, *7*, 4181, doi:10.1038/s41598-017-03739-0, 2017.

Table 1. Comparison of BC source contributions in the Arctic surface

Model and versions	Model type	Wet-deposition	Grid resolution	Meteorology	Emissions	Domain/Sites	Year/season	Major source regions/sectors
Flexpart-WRF 6.2	Lagrangian	Stohl et al. (2005)	unspecified	WRF forcast	ECLIPSE, FINN	continental Norway and Svalbard	spring 2013	Asian anthropogenic
Flexpart 6.2	Lagrangian	Stohl et al. (2005)	1° × 1°	ECMWF operational	Unspecified (BC sensitivities were calculated)	Alert, Barrow, Zeppelin	1989-2009	Northern Eurasi
Flexpart 6.2	Lagrangian	Stohl et al. (2005)	1° × 1°	ECMWF operational	ECLIPSE4(GAINS), GFED3	Arctic (north of 66°N)	2008-2010	Flaring (42%), residential (>20%
Flexpart 9.2	Lagrangian	Stohl et al. (2005)	1° × 1°	ECMWF operational	ECLIPSE5(GAINS), GFED4.1	Arctic (north of 66.7°N)	2011-2015	Residential and open burning (39%)
Flexpart 10.1	Lagrangian	Grythe et al. (2017)	1° × 1°	ECMWF operational	HTAP2, GFED3, Huang et al. (2015) for Russia flaring	Arctic (north of 66°N)	2010	Flaring (36%), open burning (18%), residentia (15%), others (31%)
GEOS-Chem 9.02	CTM	Wang et al. (2011)	2° × 2.5°	GEOS-5	HTAP2, GFED3, Huang et al. (2015) for Russia flaring	Arctic (north of 66°N)	2007-2011	Russia (62%)
GEOS-Chem	CTM	Wang et al. (2011)	2° × 2.5°	GEOS-5	Bond et al. (2004), Zhang et al. (2009), GFED3	Alert, Barrow, Zeppelin	April 2008	Asian anthropogenic (35–45%), Siberian biomas burning (46–64%
GEOS-Chem	CTM	Wang et al. (2011)	2° × 2.5°	GEOS-5	Bond et al. (2007), FLAMBE	North America Arctic	April 2008	Open fire (50%)

GEOS-Chem10.01	CTM	Wang et al. (2011)	2° × 2.5°	GEOS-5	HTAP2, ECLIPSE5, GFED4	Alert, Barrow, Zeppelin, Arctic (north of 66.5°N)	2009-2011	N-Asian anthropogenic (40–45%) in winter-spring	Xu et al. (2017)
GISS ModelE	GCM	Koch et al. (2006)	4° × 5°	Internal	Bond et al. (2004), Cooke and Wilson (1996)	Arctic (north of ~60°N)	Annual general	South Asia	Koch and Hansen (2005)

760



761

762 Figure 1. Regional separation for quantification of BC in the Arctic from (a) anthropogenic and

763 (b) open biomass burning sources.

764

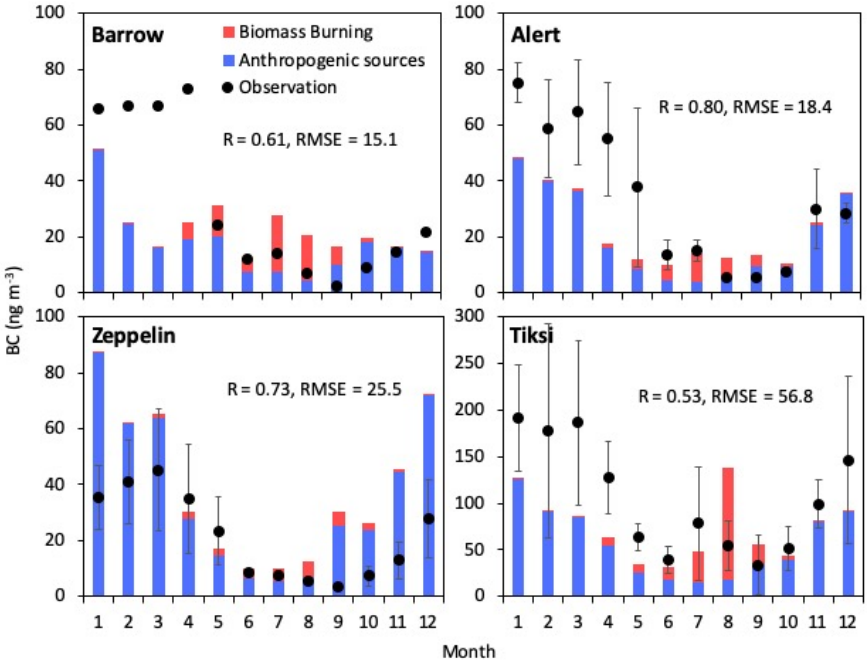


Figure 2. Observed (filled circles) and modeled (bars) seasonal variations in BC mass concentrations at Arctic sites. Contributions from anthropogenic sources (blue) and open biomass burning (red) in each month are shown. Monthly averages of observed (filled circles) and simulated (bars) BC were conducted for 2007–2011 at Alert, Canada (62.3° W, 82.5° N), and Zeppelin, Norway (11.9° E, 78.9° N), for 2009 at Barrow, USA (156.6° W, 71.3° N), and for 2010–2014 at Tiksi, Russia (128.9° E, 71.6° N). *R* and RMSE indicate correlation coefficient and root-mean-square error (ng m^{-3}), respectively.

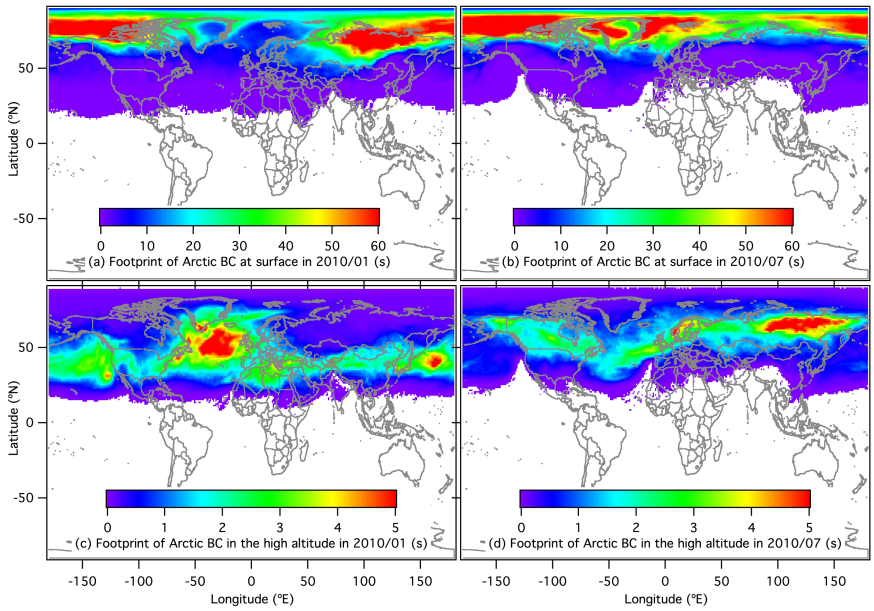


Figure 3. Footprints of Arctic BC shown as retention time(s) of (a) BC at surface (0–500 m) in January 2010, (b) BC at surface in July 2010, (c) BC at high altitudes (4750–5250 m) in January 2010, and (d) BC at high altitudes in July 2010.

Deleted: (4750–5250 m)

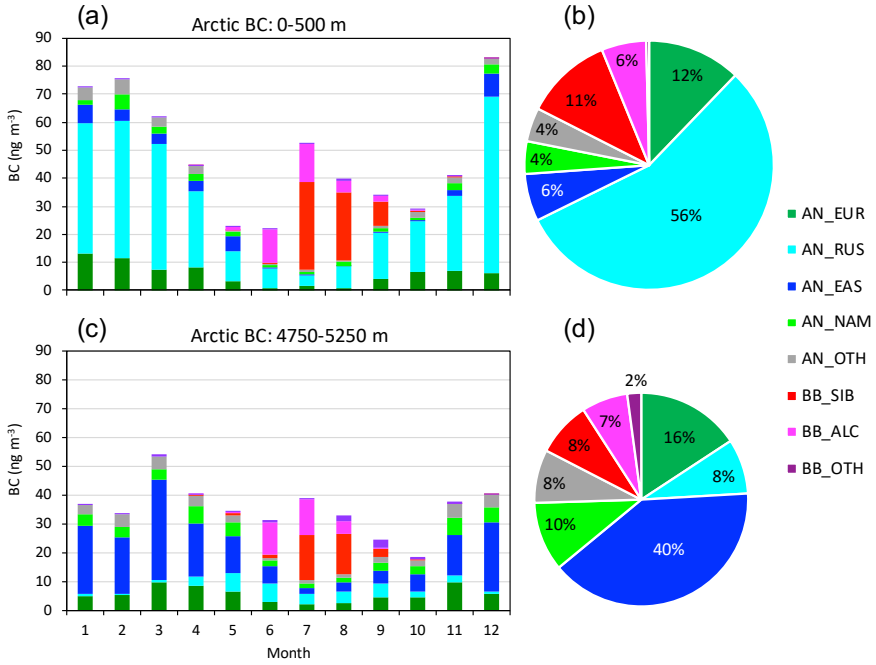
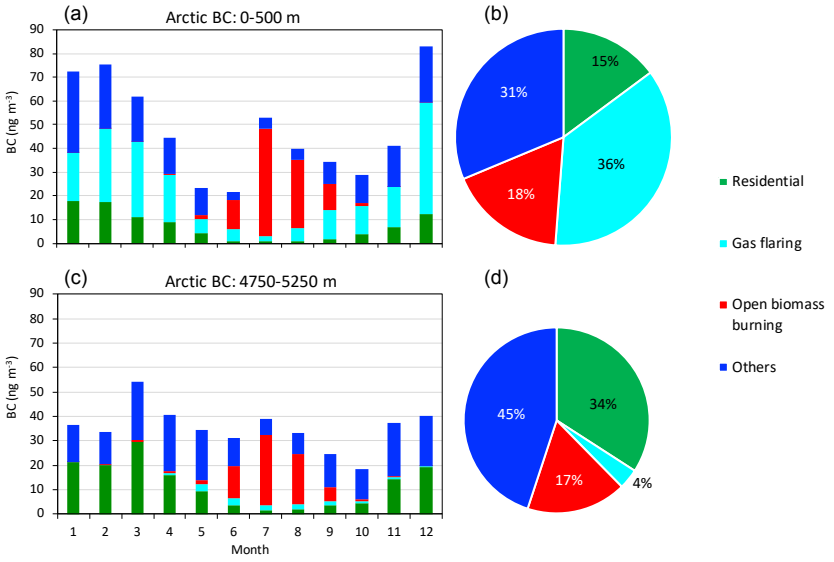


Figure 4. Contributions of anthropogenic sources (prefixed "AN_" in the legend) and open biomass burning ("BB_") from each region to (a) seasonal variations in Arctic surface BC, (b) annual mean Arctic surface BC, (c) seasonal variations in Arctic BC at high altitudes, and (d) annual mean of Arctic BC at high altitudes.



786

787 Figure 5. Sectorial contributions from residential combustion (including fossil fuel and biofuel
788 combustions), gas flaring, open biomass burning and others (energy other than gas flaring,
789 industry and transport) to (a) seasonal variations in Arctic surface BC, (b) annual mean Arctic
790 surface BC, (c) seasonal variations in Arctic BC at high altitudes, and (d) annual mean of Arctic
791 BC at high altitudes.

792

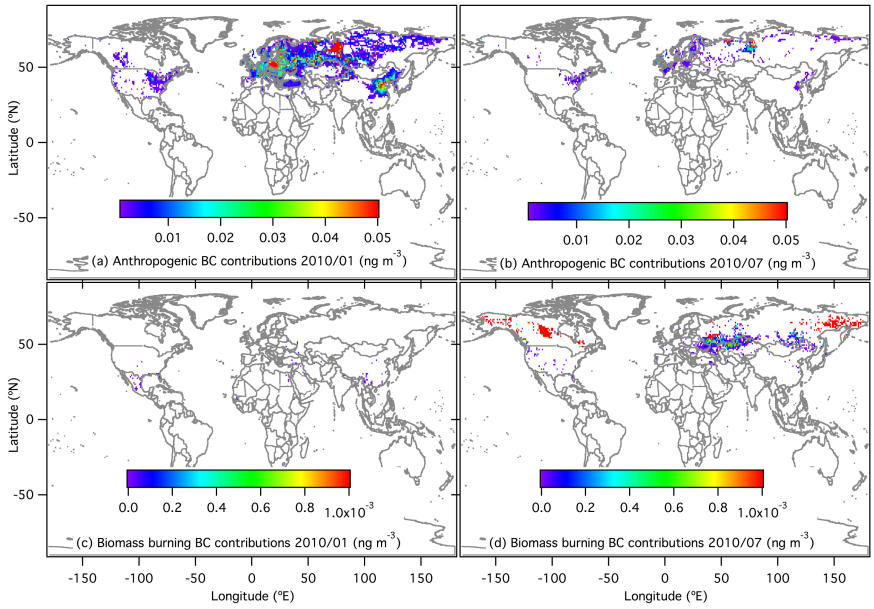


Figure 6. Spatial distributions of contributions to Arctic BC at surface for (a) anthropogenic contributions in January 2010, (b) anthropogenic contributions in July 2010, (c) open biomass burning contributions in January 2010, and (d) open biomass burning contributions in July 2010.

Deleted: (0–500 m)

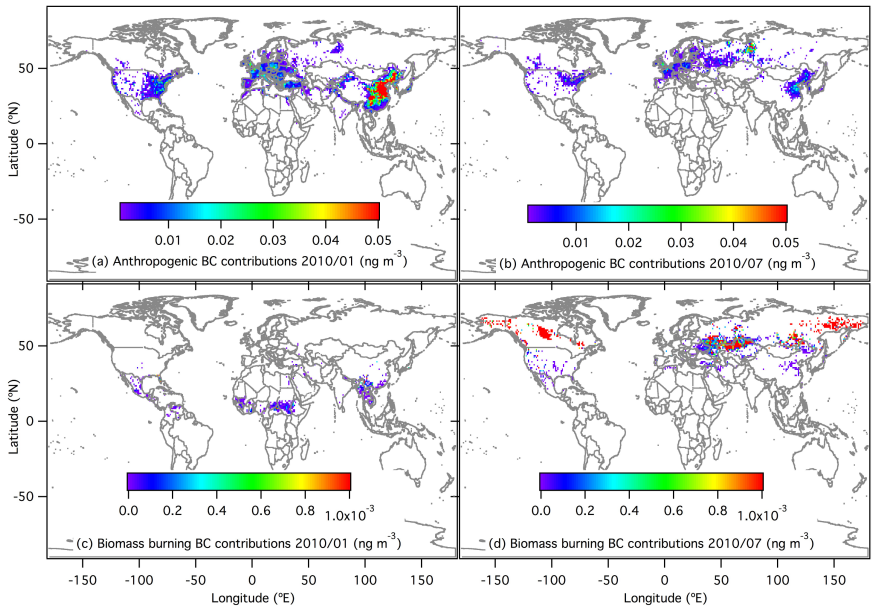


Figure 7. Spatial distributions of contributions to Arctic BC at high altitudes for (a) anthropogenic contributions in January 2010, (b) anthropogenic contributions in July 2010, (c) open biomass burning contributions in January 2010, and (d) open biomass burning contributions in July 2010.

Deleted: (4750–5250 m)