

1 **Impact of anthropogenic emissions on biogenic secondary**
2 **organic aerosol: Observation in the Pearl River Delta,**
3 **South China**

4 Yu-Qing Zhang^{1, *}, Duo-Hong Chen^{2, *}, Xiang Ding^{1, †}, Jun Li¹, Tao Zhang², Jun-Qi
5 Wang¹, Qian Cheng¹, Hao Jiang¹, Wei Song¹, Yu-Bo Ou², Peng-Lin Ye³, Gan Zhang¹,
6 Xin-Ming Wang^{1, 4}

7 1 State Key Laboratory of Organic Geochemistry and Guangdong Provincial Key Laboratory of
8 Environmental Protection and Resources Utilization, Guangzhou Institute of Geochemistry, Chinese
9 Academy of Sciences, Guangzhou, 510640, China

10 2 State Environmental Protection Key Laboratory of Regional Air Quality Monitoring, Environmental
11 Monitoring Center of Guangdong Province, Guangzhou, 510308, China

12 3 Aerodyne Research Inc., Billerica, Massachusetts 01821, United States

13 4 Center for Excellence in Regional Atmospheric Environment, Institute of Urban Environment, Chinese
14 Academy of Sciences, Xiamen, 361021, China

15 * These authors contributed equally to this work.

16 † Correspondence to: Xiang Ding (xiangd@gig.ac.cn)

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18 **Abstract.** Secondary organic aerosol (SOA) formation from biogenic precursors is affected by
19 anthropogenic emissions, which is not well understood in polluted areas. In the study, we accomplished
20 a year-round campaign at nine sites in the polluted areas located in Pearl River Delta (PRD) region during
21 2015. We measured typical biogenic SOA (BSOA) tracers from isoprene, monoterpenes, and β -
22 caryophyllene as well as major gaseous and particulate pollutants and investigated the impact of
23 anthropogenic pollutants on BSOA formation. The concentrations of BSOA tracers were in the range of
24 45.4 to 109 ng m⁻³ with the majority composed of products from monoterpenes (SOA_M, 47.2 ± 9.29 ng
25 m⁻³), followed by isoprene (SOA_I, 23.1 ± 10.8 ng m⁻³), and β -caryophyllene (SOA_C, 3.85 ± 1.75 ng m⁻³).
26 We found that atmospheric oxidants, O_x (O₃ plus NO₂), and sulfate correlated well with later-generation
27 SOA_M tracers, but not so for first-generation SOA_M products. This suggested that high O_x and sulfate
28 could promote the formation of later-generation SOA_M products, which probably led to relatively aged
29 SOA_M we observed in the PRD. For the SOA_I tracers, not only 2-methylglyceric acid (NO/NO₂-channel
30 product), but also the ratio of 2-methylglyceric acid to 2-methyltetrols (HO₂-channel products) exhibit
31 NO_x dependence, indicating the significant impact of NO_x on SOA_I formation pathways. The SOA_C tracer
32 elevated in winter at all sites and positively correlated with levoglucosan, O_x, and sulfate. Thus, the
33 unexpected increase of SOA_C in wintertime might be highly associated with the enhancement of biomass
34 burning, O₃ chemistry and sulfate components in the PRD. The BSOAs that were estimated by the SOA
35 tracer approach showed the highest concentration in fall and the lowest concentration in spring with an
36 annual average concentration of 1.68 ± 0.40 μg m⁻³. SOA_M dominated the BSOA mass all year round.
37 We also found that BSOA correlated well with sulfate and O_x. This implicated the significant effects of
38 anthropogenic pollutants on BSOA formation and highlighted that we could reduce the BSOA through
39 controlling on the anthropogenic emissions of sulfate and O_x precursors in polluted regions.

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40 **1 Introduction**

41 Secondary organic aerosols (SOA) that are produced through homogenous and heterogeneous processes
42 of volatile organic compounds (VOCs) have significant effects on global climate change and regional air
43 quality (von Schneidemesser et al., 2015). Globally, the emissions of biogenic VOCs (BVOCs) are
44 dominant over anthropogenic VOCs. Thus, biogenic SOA (BSOA) is predominant over anthropogenic
45 SOA. In the past decade, laboratorial, field, and modeling studies have demonstrated that BSOA

49 formation is highly affected by anthropogenic emissions (Zhang et al., 2015; Hoyle et al., 2011; Carlton
50 et al., 2010). Increasing NO_x shifts isoprene oxidation from the low-NO_x conditions to the high-NO_x
51 conditions (Surratt et al., 2010) and enhances nighttime SOA formation via nitrate radical oxidation of
52 monoterpenes (Xu et al., 2015). High SO₂ emission leads to abundant sulfate and acidic particles, which
53 accelerates the BSOA production by the salting-in effect and acid-catalyzed reactions (Offenberg et al.,
54 2009; Xu et al., 2016). In polluted regions, the increase of O₃ levels due to high emissions of NO_x and
55 VOCs, likely results in significant SOA formation through the ozonolysis of BVOCs (Sipilä et al., 2014;
56 Riva et al., 2017). In addition, large emission and formation of anthropogenic organic matter (OM) in
57 urban areas enhance the incorporation of BVOCs' oxidation products into the condensed phase (Donahue
58 et al., 2006). Recently, Carlton et al. (2018) found that the removal of anthropogenic emissions of NO_x,
59 SO₂, and primary OA in the CMAQ simulations could reduce BSOA by 23, 14, and 8% in summertime,
60 respectively.

61 The Pearl River Delta region (PRD) (Figure 1a) is the most developed region in China. Rapid
62 economic growth during the past three decades has resulted in large amounts of anthropogenic emissions
63 in the PRD (Lu et al., 2013). Our observation during fall-winter season in 2008 at a regional site of the
64 PRD showed that daily PM_{2.5} was as high as 150 µg m⁻³ (Ding et al., 2012). Fortunately, due to more and
65 more strict and effective pollution controls in the PRD, PM_{2.5} concentrations have significantly shrunk
66 during the last decade and met the national ambient air quality standard (NAAQS) for annual-mean PM_{2.5}
67 (35 µg m⁻³) since the year of 2015 (Figure 1b). However, O₃ and oxidant (O_x, O_x = O₃ + NO₂) are still in
68 high levels and do not decrease apparently (Figure 1b). ~~Hofzumahaus et al., (2009) observed extremely~~
69 ~~high OH concentrations in the PRD and proposed a recycling mechanism which increases the stability~~
70 ~~of OH in the air of polluted regions. All these indicate high atmospheric oxidative capacity in the PRD,~~
71 ~~since O₃, NO₂ and OH are intimately linked in atmospheric chemistry.~~ On the other hand, BVOCs
72 emissions in the PRD are expected to be high all the year in such a subtropical area (Zheng et al., 2010).
73 In the process of such a dramatic change in air pollution characteristics (e.g. PM_{2.5} and O₃), BSOA origins
74 and formation mechanisms in the PRD should be profoundly affected in the last decade. In this study,
75 year-round PM_{2.5} samples were collected at nine sites in the PRD during 2015. We investigated SOA
76 tracers from typical BVOCs (isoprene, monoterpenes, and β-caryophyllene) across the PRD for the first
77 time. We checked seasonal variations in concentrations and compositions of these BSOA tracers and
78 evaluated the impact of anthropogenic pollutants on BSOAs formation in the PRD. We also accessed the

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82 SOA origins and discussed the implication in further reducing BSOA through controlling on the
83 anthropogenic emissions.

84 **2 Experimental Section**

85 **2.1 Field Sampling**

86 Concurrent sampling was performed at 9 out of 23 sites in the Guangdong-Hong Kong-Macao regional
87 air quality monitoring network (<http://www.gdep.gov.cn/hjjce/>, Figure 1a), including three urban sites in
88 Zhaoqing (ZQ), Guangzhou (GZ) and Dongguan (DG), two suburban sites in Nansha (NS) and Zhuhai
89 (ZH), and four rural sites in Tianhu (TH), Boluo (BL), Heshan (HS) and Taishan (TS).

90 At each site, 24-hr sampling was conducted every six days from January to December in 2015 using
91 a PM_{2.5} sampler equipped with quartz filters (8 × 10 inches) at a flow rate of 1.1 m³ min⁻¹. Additionally,
92 field blanks were collected monthly at all sites. Blank filters were covered with aluminum foil and baked
93 at 500 °C for 12 hrs and stored in a container with silica gel. After sampling, the filter samples were
94 stored at -20 °C.

95 In this study, the filters collected in January, April, July and October 2015 were selected to represent
96 winter, spring, summer, and fall samples, respectively. A total of 170 field samples (4-5 samples for each
97 season at each site) were analyzed in the current study.

98 **2.2 Chemical Analysis**

99 For each filter, organic carbon (OC) and elemental carbon (EC) were measured by an OC-EC aerosol
100 analyzer (Sunset Laboratory Inc.). Water-soluble ions were analyzed by ion chromatography (Metrohm).
101 All these species are major components in PM_{2.5} (see Figure 2). Meteorological parameters (temperature
102 and relative humidity) and gaseous pollutants (SO₂, CO, NO₂, NO, and O₃) at each site were recorded
103 hourly. We further calculated the daily averages to probe the potential influence of air pollutants on
104 BSOA formation.

105 For BSOA tracer analysis, detailed information of the processes is described in the previous
106 literatures (Shen et al., 2015; Ding et al., 2012). Isotope-labeled standard mixtures, including dodecanoic
107 acid-d₂₃, hexadecanoic acid-d₃₁, docosanoic acid-d₄₃ and levoglucosan-¹³C₆ were added into each sample
108 as internal standards. Then, samples were extracted by sonication with the mixed solvents of dichloride
109 methane (DCM)/hexane (1:1, v/v) and DCM/methanol (1:1, v/v), sequentially. The extraction solutions

110 of each sample were combined, filtered, and concentrated to ~2 mL. Each concentrated sample was split
111 into two parts for silylation and methylation, respectively.

112 We analyzed fourteen BSOA tracers in the derivatized samples using GC/MSD (Agilent
113 7890/5975C). The isoprene-derived SOA (SOA_I) tracers were composed of 2-methyltetrols (2-MTLs, 2-
114 methylthreitol and 2-methylerythritol) (Claeys et al., 2004a), 2-methylglyceric acid (2-MGA) (Claeys et
115 al., 2004b), 3-MeTHF-3,4-diols (cis-3-methyltetrahydrofuran-3,4-diol and trans-3-
116 methyltetrahydrofuran-3,4-diol) (Lin et al., 2012) and C₅-alkene triols (cis-2-methyl-1,3,4-trihydroxy-1-
117 butene, trans-2-methyl-1,3,4-trihydroxy-1-butene and 3-methyl-2,3,4-trihydroxy-1-butene) (Wang et al.,
118 2005). The monoterpenes-derived SOA (SOA_M) tracers included 3-hydroxy-4,4-dimethylglutaric acid
119 (HDMGA), 3-hydroxyglutaric acid (HGA) (Claeys et al., 2007), pinic acid (PA), cis-pinonic acid (PNA)
120 (Christoffersen et al., 1998), and 3-methyl-1,2,3-butanetricarboxylic acid (MBTCA) (Szmigielski et al.,
121 2007). The β-caryophyllene-derived SOA (SOA_C) tracer was β-caryophyllenic acid (CA) (Jaoui et al.,
122 2007). Due to the lack of authentic standards, surrogate standards were used to quantify BSOA tracers
123 except PNA. Specifically, erythritol, PNA and octadecanoic acid were used for the quantification of SOA_I
124 tracers (Ding et al., 2008), other SOA_M tracers (Ding et al., 2014) and CA (Ding et al., 2011), respectively.
125 The method detection limits (MDLs) for erythritol, PNA and octadecanoic acid were 0.01, 0.02, and 0.02
126 ng m⁻³, respectively. Table S1 summarizes BSOA data at each site in the PRD.

127 2.3 Quality Assurance / Quality Control

128 These target BSOA tracers were not detected or lower than MDLs in the field blanks. The results of
129 spiked samples (erythritol, PNA and octadecanoic acid spiked in pre-baked quartz filters) indicated that
130 the recoveries were 65 ± 14 % for erythritol, 101 ± 3 % for PNA, and 83 ± 7 % for octadecanoic acid.
131 The results of paired duplicate samples indicated that all the relative differences for target BSOA tracers
132 were lower than 15%.

133 It should be noted that the application of surrogate quantification introduces additional errors to the
134 results. Based on the empirical approach to calculate uncertainties from surrogate quantification (Stone
135 et al., 2012), we estimated the errors in analyte measurement which were propagated from the
136 uncertainties in field blanks, spike recoveries, repeatability and surrogate quantification. As Table S2
137 showed, the estimated uncertainties in the tracers' measurement ranged from 15% (PNA) to 157% (CA).

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140 3 Results and Discussion

141 3.1 PM_{2.5} and gaseous pollutants

142 Figure 2 presents spatial and seasonal variations of PM_{2.5} and its major components. Although annual-
143 mean PM_{2.5} ($34.8 \pm 6.1 \mu\text{g m}^{-3}$) in the PRD met the NAAQS value of $35 \mu\text{g m}^{-3}$, PM_{2.5} at the urban sites
144 (ZQ, GZ and DG) all exceeded the NAAQS value. The rural TH site in the northern part of PRD
145 witnessed the lowest concentration of PM_{2.5} ($25.0 \mu\text{g m}^{-3}$) among the nine sites. PM_{2.5} levels were highest
146 in winter (on average $60.1 \pm 21.6 \mu\text{g m}^{-3}$) and lowest in summer (on average $22.8 \pm 3.3 \mu\text{g m}^{-3}$).
147 Carbonaceous aerosols and water-soluble ions together explained $98 \pm 11 \%$ of PM_{2.5} masses. OM
148 (OC \times 1.6) was the most abundant component in PM_{2.5}, followed by sulfate, ammonium, nitrate and EC.
149 Similar to PM_{2.5}, the five major components all increased in winter and fall (Figure S1), suggesting severe
150 PM_{2.5} pollution during fall-winter season in the PRD.

151 In the gas phase, SO₂, CO, NO₂ and NO_x presented similar seasonal trends as PM_{2.5}, i.e. higher
152 levels occurred during fall and winter and lower concentrations during spring and summer (Figure 3 a-
153 d). Annual-mean SO₂ and NO₂ in the PRD both met the NAAQS values of $60 \mu\text{g m}^{-3}$ and $40 \mu\text{g m}^{-3}$,
154 respectively (Figure 3a and 3c). As a typical secondary pollutant, O₃ was highest in summer (Figure 3e),
155 probably because of the strong photo-chemistry. Due to the compromise of opposite seasonal trends of
156 O₃ and NO₂, O_x showed less seasonal variation (Figure 3f) compared with other gaseous pollutants. And
157 annual-mean O_x reached $96.1 \pm 14.9 \mu\text{g m}^{-3}$. These indicated significant O₃ pollution all the year in the
158 PRD.

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159 3.2 Spatial distribution and seasonal variation of SOA tracers

160 The total concentrations of BSOA tracers ranged from 45.4 to 109 ng m^{-3} among the nine sites. SOA_M
161 tracers ($47.2 \pm 9.29 \text{ ng m}^{-3}$) represented predominance, followed by SOA_I tracers ($23.1 \pm 10.8 \text{ ng m}^{-3}$),
162 and SOA_C tracer ($3.85 \pm 1.75 \text{ ng m}^{-3}$).

163 3.2.1 Monoterpenes-derived SOA tracers

164 Annual averages of total SOA_M tracers at the nine sites were in the range of 26.5 to 57.4 ng m^{-3} (Table
165 S1). Figure 4 and Figure S2a show the spatial distribution of SOA_M tracers and monoterpene emissions
166 in the PRD (Zheng et al., 2010). The highest concentration of SOA_M tracers was observed at the rural
167 TH site where monoterpene emissions were high. Figure 4 also presents seasonal variations of SOA_M

170 tracers. At most sites, high levels occurred in summer and fall. Monoterpene emission rates are
171 influenced by temperature and solar radiation (Guenther et al., 2012). Thus, high temperature and
172 intensive solar radiation during summer and fall in the PRD (Zheng et al., 2010) could stimulate
173 monoterpene emissions and then the SOA_M formation.

174 Among the five SOA_M tracers, HGA ($20.1 \pm 4.28 \text{ ng m}^{-3}$) showed the highest concentration,
175 followed by HDMGA ($14.7 \pm 2.93 \text{ ng m}^{-3}$), MBTCA ($7.63 \pm 1.49 \text{ ng m}^{-3}$), PNA ($3.75 \pm 2.72 \text{ ng m}^{-3}$) and
176 PA ($1.01 \pm 0.48 \text{ ng m}^{-3}$). SOA_M formation undergoes multi-generation reactions. The first-generation
177 SOA_M (SOA_{M,F}) products, PNA and PA, can be further oxidized and form the later-generation (SOA_{M,L})
178 products, e.g. MBTCA (Müller et al., 2012). Thus, the (PNA+PA) / MBTCA ratio has been used to probe
179 SOA_M aging (Haque et al., 2016; Ding et al., 2014). The (PNA+PA) / MBTCA ratios in chamber-
180 generated α -pinene SOA samples were reported in the range of 1.51 to 5.91 depending on different
181 oxidation conditions (Offenberg et al., 2007; Eddingsaas et al., 2012). In this study, the median values of
182 (PNA+PA) / MBTCA varied from 0.27 at ZH to 1.67 at TH. The ratios observed in this study were
183 consistent with our previous observations at the regional site, Wanqingsha (WQS) in the PRD (Ding et
184 al., 2012), but lower than those in the fresh α -pinene SOA samples from chamber experiments (Figure
185 S3), indicating relatively aged SOA_M in the air of PRD.

186 Moreover, the levels of SOA_{M,L} tracers (HGA + HDMGA + MBTCA) were much higher than those
187 of SOA_{M,F} tracers (PNA + PA), with mean mass fractions of SOA_{M,L} tracers reaching 86% (Figure 4).
188 Mass fractions of SOA_{M,F} tracers decreased in the summer samples (Figure 4), probably resulting from
189 strong photo-chemistry and more intensive further oxidation during summer. High abundances of
190 SOA_{M,L} tracers in the PRD were different from our year-round observations at 12 sites across China
191 (Ding et al., 2016b). In that study, the (PNA+PA) / MBTCA ratio suggested generally fresh SOA_M (Figure
192 S3) and SOA_{M,F} tracers were the majority. Thus, we see more aged SOA_M in the PRD.

193 As Figure 5 a-b and Figure S4,S5 showed, the SOA_{M,F} tracers did not show good correlations with
194 O_x at most sites, while the SOA_{M,L} tracers exhibited significant O_x dependence. When O_x is high, strong
195 photo-oxidation of PNA and PA could reduce their concentrations and promote the formation of SOA_{M,L}
196 tracers (Müller et al., 2012). Thus, the levels of SOA_{M,L} tracers would increase with increasing O_x but
197 not so for SOA_{M,F} tracers. On the other hand, sulfate is a key species in particles that determines aerosol
198 liquid water amount, aerosol acidity and particle surface area (Xu et al., 2015, 2016). Thus, the increase
199 of sulfate could promote aqueous and heterogeneous reactions. In this study, the SOA_{M,F} tracers poorly

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213 correlated with sulfate (Figure 5c), while the SOA_{M,L} tracers positively correlated with sulfate at all the
214 9 sites (Figure 5d). At each site the SOA_{M,L} tracers exhibited more sulfate dependence than the SOA_{M,F}
215 tracers (Figure S5). This suggested that sulfate also played a critical role in forming SOA_{M,L} tracers
216 through the particle-phase reactions. Besides the gas-phase OH oxidation (Müller et al., 2012), the
217 heterogeneous OH oxidation of pinonic acid could also produce SOA_{M,L} tracers (Lai et al. 2015).
218 Aljawhary et al., (2016) reported the kinetics and mechanism of pinonic acid oxidation in acidic solutions
219 and found that the molar yields of MBTCA through the aqueous-phase reactions were similar to those in
220 the gas-phase oxidation. Here, we conclude that high concentrations of O_x and sulfate could stimulate
221 SOA_{M,L} tracers' production and thereby lead to aged SOA_M in the PRD.

222 3.2.2 Isoprene-derived SOA tracers

223 Annual averages of total SOA_I tracers at the nine sites were in the range of 10.8 to 49.3 ng m⁻³ (Table
224 S1). Figure 6 and Figure S2b show the spatial distribution of SOA_I tracers and isoprene emissions in the
225 PRD (Zheng et al., 2010), respectively. The highest concentration occurred at ZQ where the emissions
226 were high. Figure 6 also presents seasonal variations of SOA_I tracers at the nine sites. High levels
227 occurred in summer and fall. Similar to monoterpenes, the emission rate of isoprene is influenced by
228 temperature and solar radiation (Guenther et al., 2012), which are expected to be higher in summer and
229 fall in the PRD (Zheng et al., 2010). Among these SOA_I tracers, 2-MTLs (14.2 ± 5.61 ng m⁻³) were the
230 most abundant products, followed by C₅-alkene triols (6.81 ± 5.05 ng m⁻³), 2-MGA (1.99 ± 0.72 ng m⁻³)
231 and 3-MeTHF-3,4-diols (0.19 ± 0.08 ng m⁻³).

232 SOA_I formation is highly affected by NO_x (Surratt et al., 2010). Under the low-NO_x or NO_x free
233 conditions, isoprene is oxidized by the OH and HO₂ radicals through the HO₂-channel which generates
234 a hydroxy hydroperoxide (ISOPOOH) and then forms epoxydiols (IEPOX) (Paulot et al., 2009). Reactive
235 uptake of IEPOX on acidic particles eventually produces 2-MTLs, C₅-alkene triols, 3-MeTHF-3,4-diols,
236 2-MTLs-organosulfates and oligomers (Lin et al., 2012). Under the high-NO_x conditions, isoprene
237 undergoes oxidation by NO_x through the NO/NO₂-channel and generates methacrolein (MACR) and then
238 forms peroxyethylacrylic nitric anhydride (MPAN). Further oxidation of MPAN by the OH radical
239 produces hydroxymethyl-methyl- α -lactone (HMML) and/or methacrylic acid epoxide (MAE). HMML
240 and MAE are the direct precursors to 2-MGA, 2-MGA-organosulfate and its corresponding oligomers
241 (Nguyen et al., 2015). As Figure 6 showed, the concentrations of HO₂-channel tracers (2-MTLs + C₅-

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248 alkene triols + 3-MeTHF-3,4-diols) were much higher than those of the NO/NO₂-channel product (2-
249 MGA) at all the nine sites. The dominance of HO₂-channel products was also observed at another
250 regional site in the PRD (WQS)₁ (He et al., 2018).

251 Figure 6 also shows seasonal trends of the 2-MGA to 2-MTLs ratio (2-MGA/2-MTLs) which is
252 often applied to probe the influence of NO_x on the formation of SOA₁ (Ding et al., 2013; Ding et al.,
253 2016a; Pye et al., 2013). The ratios were highest in wintertime and lowest in summertime, which were
254 consistent with the seasonal trend of NO_x during our campaign (Figure 3d). As Table 1 showed, 2-MGA
255 positively correlated with NO₂, probably due to the enhanced formation of MPAN from
256 peroxyacetyl (PMA) radical reacted with NO₂ (Worton et al., 2013; Chan et al., 2010). Previous
257 laboratory studies showed that increasing NO₂/NO ratio could promote the formation of 2-MGA and its
258 corresponding oligoesters (Chan et al., 2010; Surratt et al., 2010). However, we did not see a significant
259 correlation between 2-MGA and NO₂/NO ratio in the PRD. Instead, the 2-MGA/2-MTLs ratio correlated
260 well with NO, NO₂ and NO₂/NO ratio (Table 1). Increasing NO limits the formation of ISOPOOH but
261 prefers the production of MACR, and increasing NO₂ enhances MPAN formation. Thus, it is expected
262 that the 2-MGA/2-MTLs ratio shows stronger NO_x dependence than 2-MGA. These findings demonstrate
263 the significant impact of NO_x on SOA₁ formation pathways in the atmosphere. We also checked the
264 correlations of SOA₁ tracers with O₃ and sulfate (Figure S6). The NO/NO₂-channel product exhibited
265 more O₃ and sulfate dependence than HO₂-channel products.

266 Recent studies indicated that isoprene ozonolysis might play a role in SOA₁ formation in the ambient
267 air. Riva et al. (2016) found that isoprene ozonolysis with acidic particles could produce substantial 2-
268 MTLs but not so for C₅-alkene triols and 3-MeTHF-3,4-diols. Li et al. (2018) observed a positive
269 correlation between 2-MTLs and O₃ in the North China Plain. In the PRD, we also saw weak but
270 significant correlations of 2-MTLs with O₃ (Table S3). However, 3-MeTHF-3,4-diols and C₅-alkene
271 triols were detected in all samples and 2-MTLs, C₅-alkene triols and 3-MeTHF-3,4-diols correlated well
272 with each other (Table S4), which was apparently different from those reported by Riva et al. (2016).
273 Moreover, the ratios of 2-MTLs isomers in the PRD samples (2.00–2.85) were much lower than those
274 (10–22, Figure S7) reported in the SOA from isoprene ozonolysis (Riva et al., 2016). Furthermore,
275 isoprene oxidation by the OH radical is much faster than that by ozone under the polluted PRD conditions
276 (Table S5). And IEPOX yields through the ISOPOOH oxidation by the OH radical are more than 75%
277 in the atmosphere (St. Clair et al., 2016). Thus, isoprene ozonolysis might be not the major formation

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281 pathway of SOA₁, even though annual-mean O₃ level reaching 67.7 μg m⁻³ in the PRD (Table S1).

282 Previous studies found that thermal decomposition of low volatility organics in IEPOX-derived
283 SOA could produce SOA₁ tracers, e.g. 2-MTLs, C₅-alkene triols and 3-MeTHF-3,4-diols (Lopez-Hilfiker
284 et al., 2016, Watanabe et al., 2018). This means that these tracers detected by GC-MSD might be
285 generated from thermal decomposition of IEPOX-derived SOA. As estimated by Cui et al (2018), 14.7-
286 42.8% of C₅-alkene triols, 11.1% of 2-MTLs and approximately all 3-MeTHF-3,4-diols measured by
287 GC/MSD could be attributed to the thermal degradation of 2-MTLs-derived organosulfates (MTL-OSs).
288 We also measured MTL-OSs in two samples at HS and TS sites, respectively (Table S6) using the widely
289 used LC-MS approach (He et al., 2014, 2018). Assuming that all MTL-OSs decomposed to these tracers,
290 the thermal decomposition of MTL-OSs would account for 15.1-31.6% of C₅-alkene triols, 6.0-10.0% of
291 2-MTLs and all 3-MeTHF-3,4-diols measured by GC/MSD. Thus, C₅-alkene triols and 2-MTLs are
292 major from isoprene oxidation rather than thermal decomposition of MTL-OSs, while 3-MeTHF-3,4-
293 diols are only in trace amount in the air and might be produced largely from thermal degradation.

294 Moreover, we see significant variations in SOA₁ tracer compositions in the PRD. For instant, C₅-
295 alkene triols have three isomers. If these tracers were mainly generated from a thermal process, their
296 compositions should be similar in different samples. In fact, the relative abundances of three C₅-alkene
297 triol isomers significantly changed from site to site (Figure 7) and season to season (Figure S8), and their
298 compositions in the PRD were different from those measured in the chamber samples (Lin et al., 2012).
299 In addition, the slopes of linear correlations among these IEPOX-derived SOA tracers also varied from
300 site to site (Figure S9). Coupled with the seasonal trend of 2-MGA/2-MTLs ratios, the apparent variations
301 in SOA₁ tracer compositions demonstrate that these SOA₁ tracers are mainly formed through different
302 pathways in the ambient atmosphere, although part of them might arise from the thermal decomposition
303 of different dimers/OSs and the parent dimers/OSs varies with sites and seasons.

304 **3.2.3 Sesquiterpene-derived SOA tracer**

305 Annual averages of CA at the nine sites ranged from 1.82 to 7.07 ng m⁻³. The levels of CA at the inland
306 sites (e.g. GZ, ZQ, and TH) were higher than those at the coastal sites (ZH and NS, Figure 8). Since
307 sesquiterpenes are typical BVOCs, it is unexpected that the concentrations of CA were highest during
308 winter in the PRD (Figure 8). Interestingly, seasonal trend of CA was consistent with that of the biomass
309 burning (BB) tracer, levoglucosan (Figure 8). And CA correlated well with levoglucosan at eight sites in

删除的内容: A previous study found that thermal decomposition of low volatility organics in IEPOX-derived SOA could produce SOA₁ tracers, e.g. 2-MTLs, C₅-alkene triols and 3-MeTHF-3,4-diols (Lopez-Hilfiker et al., 2016). This means that these tracers detected by GC-MSD might be generated from thermal decomposition of IEPOX-derived SOA. If these tracers were mainly generated from such a thermal process, their compositions would be similar in different samples. To verify this possibility, we presented chemical compositions of three C₅-alkene triol isomers at the nine sites in ternary plots. The relative abundances of three isomers significantly changed from site to site (Figure 7) and season to season (Figure S5), and were different from those measured in the chamber samples (Lin et al., 2012). Moreover, the slopes of linear correlations among these IEPOX-derived SOA tracers also varied from site to site (Figure S6). Coupled with the seasonal trend of 2-MGA/2-MTLs ratios, the observed variations in SOA₁ tracers compositions demonstrated that the SOA₁ tracers were mainly formed through different pathways in the real atmosphere rather than thermal decomposition. .

331 the PRD (Figure 9a). Sesquiterpenes are stored in plant tissues partly to protect the plants from insects
332 and pathogens (Keeling and Bohlmann, 2006). BB can not only stimulate sesquiterpene emissions
333 (Ciccioli et al., 2014) but also substantially alter the SOA formation and yields (Mentel et al., 2013).
334 Emissions inventories in the PRD showed that the BB emissions were enhanced during winter (He et al.,
335 2011). These suggested that the unexpected increase of SOA_C in wintertime could be highly associated
336 with BB emissions in the PRD.

337 Besides the impact of BB, we also found positive correlations of CA with O_x (Figure 9b) and sulfate
338 (Figure 9c). The oxidation of β-caryophyllene by the OH radical and O₃ is very rapid. Under typical
339 oxidation conditions in the air of PRD, the lifetimes of β-caryophyllene are only several minutes (Table
340 S5). Once emitted from vegetation or biomass burning, β-caryophyllene will react rapidly and form
341 CA immediately. This partly explains the positive correlations between CA and levoglucosan in the PRD.
342 The unexpected high levels of CA in the winter indicated that biomass burning could be an important
343 source of SOA_C in the PRD, especially in wintertime. In addition, the increase of sulfate could raise
344 aerosol acidity and thereby promote aqueous and heterogeneous reactions to form SOA_C. In the PRD,
345 both O_x (Figure 3f) and sulfate (Figure S1) increased during winter, which could promote SOA_C
346 formation. Here, we conclude that the enhancement of BB emissions as well as the increase of O_x and
347 sulfate in wintertime together led to high SOA_C production during winter in the PRD.

348 3.3 Source apportionment and atmospheric implications

349 We further attributed BSOA by the SOA-tracer approach which was first developed by Kleindienst et al.
350 (2007). This method has applied to SOA apportionment at multiple sites across the United States
351 (Lewandowski et al., 2013) and China (Ding et al., 2016b), and over global oceans from Arctic to
352 Antarctic (Hu et al., 2013). Details of the SOA-tracer method and its application in this study as well as
353 the uncertainty of estimating procedure are described in Text S1. Table S1 lists the results of estimated
354 SOA from different BVOCs.

355 Figure 10a exhibits the spatial distribution of BSOA (SOA_M + SOA_I + SOA_C). Annual average at
356 the nine sites ranged from 0.97 μg m⁻³ (NS) to 2.19 μg m⁻³ (ZQ), accounting for 9-15% of OM. SOA_M
357 was the largest BSOA contributor with an average contribution of 64 ± 7 %, followed by SOA_C (21 ±
358 6 %), and SOA_I (14 ± 4 %). Figure 10b presents seasonal variation of BSOA. The levels were highest in
359 fall (2.35 ± 0.95 μg m⁻³) and lowest in spring (1.06 ± 0.42 μg m⁻³). SOA_M contributions ranged from 57%

删除的内容: The oxidation of β-caryophyllene by the OH radical and O₃ is very rapid with the lifetimes less than 10 min under typical conditions in the air of PRD (Table S5). The increase of sulfate could not only raise aerosol acidity but also enhance the salting-in effect (Xu et al., 2015). In the PRD, both O_x (Figure 3f) and sulfate (Figure S1) increased during winter, which could promote SOA_C formation.

367 in winter to 68% in spring. The shares of SOA₁ were only 5% in winter and reached up to 22% in summer.
368 The contributions of SOA_C increased to 40% in wintertime.

369 It is interesting to note that SOA_M, SOA₁ and SOA_C all positively correlated with sulfate and O_x in
370 the PRD (Table 2). Since anthropogenic emissions can enhance BSOA formation (Hoyle et al., 2011),
371 the reduction of anthropogenic emissions indeed lowers BSOA production (Carlton et al., 2018). As the
372 oxidation product of SO₂, sulfate is a key species in particles that determines aerosol acidity and surface
373 areas (Xu et al., 2015, 2016) which could promotes BSOA formation through the acid-catalyzed
374 heterogeneous reactions. Recent study found that SO₂ could directly reaction with organic peroxides of
375 monoterpene ozonolysis and form substantial organosulfates (Ye et al., 2018). Thus, the decrease of SO₂
376 emission indeed reduces SO₂ and sulfate in the ambient air, which hereby leads to less acidic particles
377 and reduces the BSOA production. For O_x, the increase of O₃ likely results in significant SOA formation
378 through the BVOCs ozonolysis (Sipilä et al., 2014; Riva et al., 2017). Hence, the decrease of O_x resulting
379 from the control of VOCs and NO_x emissions could reduce BSOA formation through O₃ chemistry. Based
380 on the observed sulfate and O_x dependence of BSOA in this study, the reduction of 1 μg m⁻³ in sulfate
381 and O_x in the air of PRD could lower BSOA levels by 0.17 and 0.02 μg m⁻³, respectively. If both
382 concentrations decline by 50%, the reduction of O_x is more efficient than sulfate in reducing BSOA in
383 the PRD (Table 2).

384 We further compared the results in 2015 with those during fall-winter season in 2008 at WQS (Ding
385 et al., 2012). We found that all BSOA species positively correlated with sulfate but exhibited no O_x
386 dependence (Table S7). Thus, in 2008 BSOA formation was largely influenced by sulfate, probably due
387 to high sulfate levels then (as high as 46.8 μg m⁻³). Owing to strict control of SO₂ emissions (Wang et al.,
388 2013), ambient SO₂ significantly shrank over the PRD (Figure 1b). Our long-term observation during
389 fall-winter season at WQS also witnessed a decreasing trend of sulfate from 2007 to 2016 (Figure S10).
390 However, O_x levels did not decrease during the past decade (Figure 1b) and O_x concentrations were much
391 higher than sulfate in 2015 in the PRD (96.1 ± 14.9 μg m⁻³ vs. 8.44 ± 1.09 μg m⁻³ on average). All these
392 underline the importance of O_x in BSOA formation currently in the PRD. At present, short-term despiking
393 and long-term attainment of O₃ concentrations are challenges for air pollution control in the PRD (Ou et
394 al., 2016). Thus, lowering O_x is critical to improve air quality in the PRD. Our results highlight the
395 importance of future reduction in anthropogenic pollutant emissions (e.g. SO₂ and O_x precursors) for
396 considerably reducing the BSOA burden in polluted regions.

删除的内容: Anthropogenic emissions significantly influence BSOA formation (Carlton et al., 2018). The observed sulfate and O_x dependence of BSOA in the air of PRD indicates that the reduction of 1 μg m⁻³ in sulfate and O_x could lower BSOA levels by 0.17 and 0.02 μg m⁻³, respectively. If both concentrations decline by 50%, the reduction of O_x is more efficient than sulfate in reducing BSOA in the PRD (Table 2). Due to strict control of SO₂ emissions (Wang et al., 2013), ambient SO₂ significantly shrank over the PRD (Figure 1b). However, O_x levels did not decrease during the past decade (Figure 1b) and O_x concentrations were much higher than sulfate in the PRD (96.1 ± 14.9 μg m⁻³ vs. 8.44 ± 1.09 μg m⁻³ on average).

410 **Code/Data availability**

411 The experimental data in this study are available upon request to the corresponding author by email.

412 **Author Contribution**

413 Xiang Ding, Duo-Hong Chen and Jun Li conceived the project and designed the study. Yu-Qing Zhang
414 and Duo-Hong Chen performed the data analysis and wrote the manuscript. Duo-Hong Chen, Tao Zhang
415 and Yu-Bo Ou arranged the sample collection and assisted with the data analysis. Jun-Qi Wang, Qian
416 Cheng and Hao Jiang analyzed the samples. Xiang Ding, Peng-Lin Ye, Wei Song, Gan Zhang and Xin-
417 Ming Wang performed data interpretation and edited the manuscript. All authors contributed to the final
418 manuscript development.

419 **Competing interests**

420 The authors declare that they have no conflict of interest.

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425 [would like to thank Prof. Sasho Gligorovski for his helpful suggestion on the discussion of atmospheric](#)
426 [oxidation process.](#) The data of gaseous pollutants, major components in PM_{2.5} and BSOA tracers can be
427 found in supporting information.

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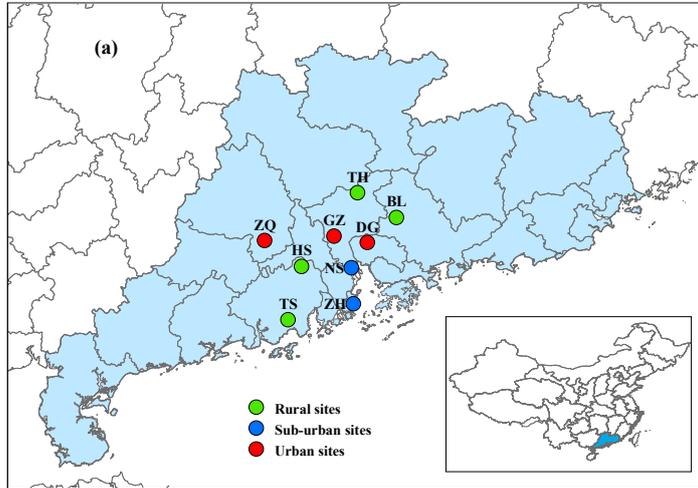
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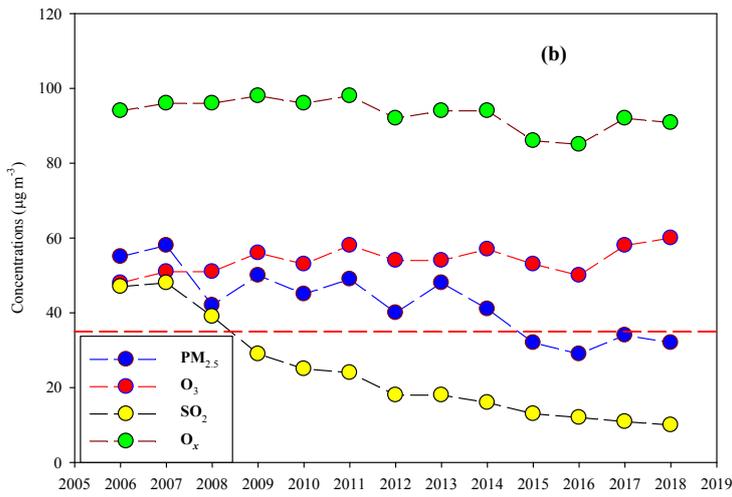
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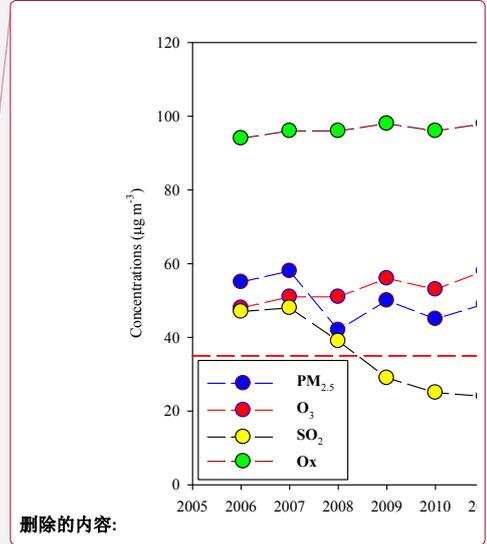
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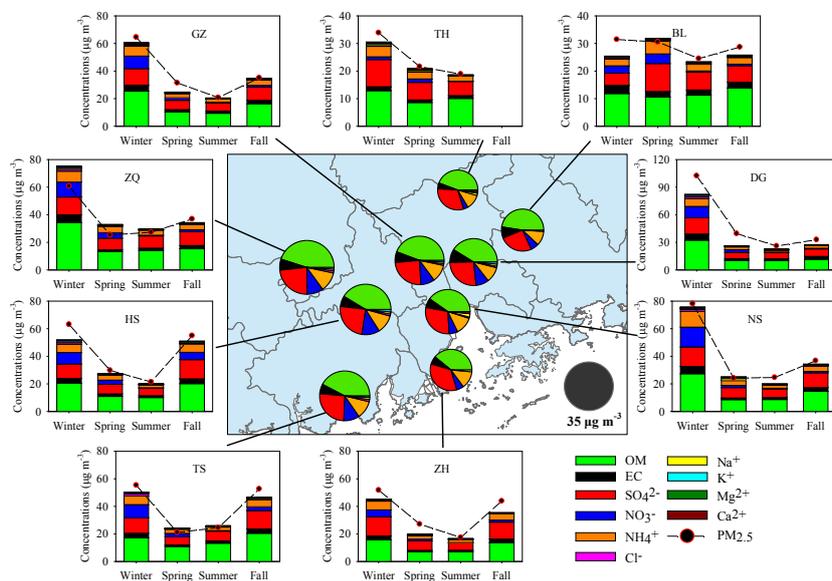
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661 Figure 1 Sampling sites in the PRD (a) and long-term trends of annual-mean $PM_{2.5}$, O_3 , SO_2 and O_x recorded by the
 662 Guangdong-Hong Kong-Macao regional air quality monitoring network (<http://www.gdep.gov.cn/hjice/>) (b). The
 663 red dash line indicates the NAAQS for annual-mean $PM_{2.5}$ concentrations ($35 \mu g m^{-3}$).

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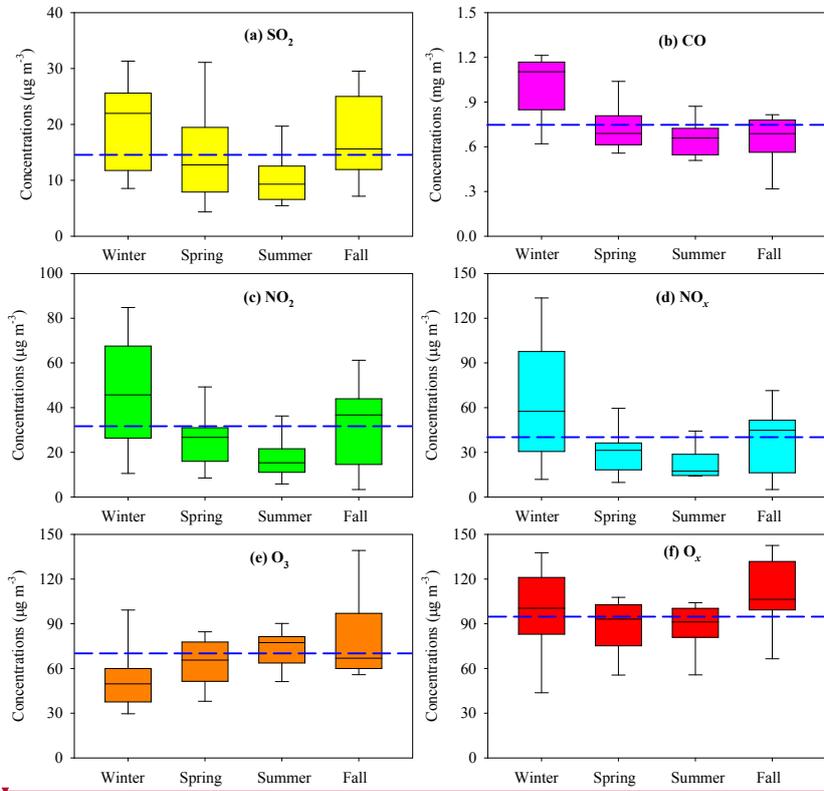


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666
 667 Figure 2 Major components in PM_{2.5} and their seasonal variation at 9 sites. The pie charts in the central figure
 668 represent the annual average of major components. High levels of PM_{2.5} and major components were observed in
 669 wintertime.

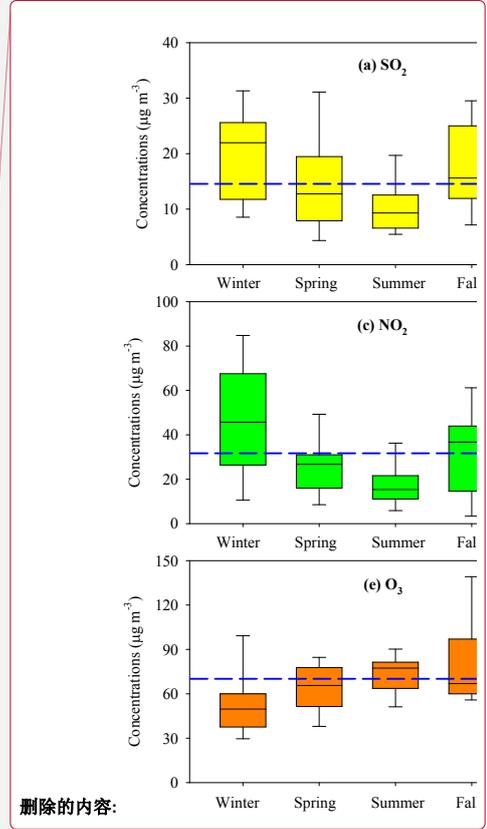
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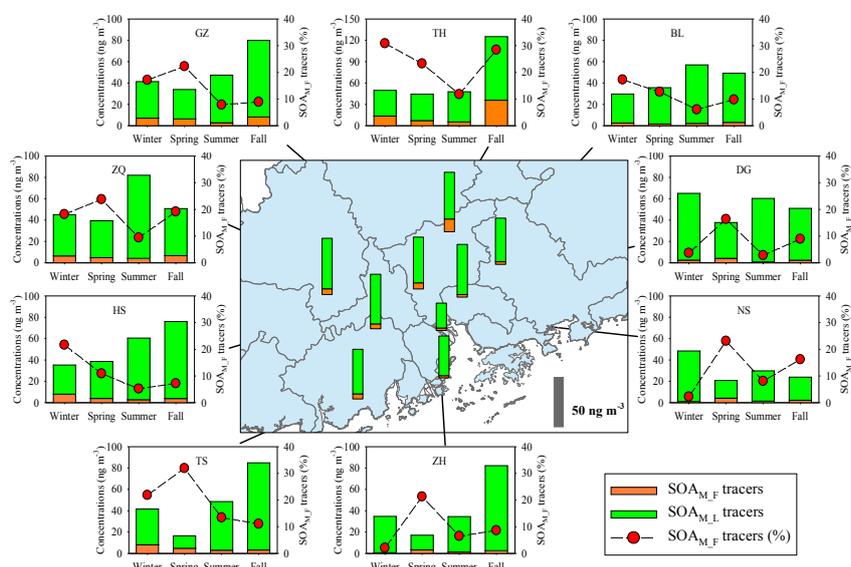


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672 Figure 3 Seasonal variation of gaseous pollutants in the PRD. Box with error bars represent 10th, 25th, 75th, 90th
 673 percentiles for each pollutant. The line in each box represents the median value. Blue dash lines indicate annual
 674 average concentrations of SO_2 ($14.9 \mu\text{g m}^{-3}$), CO (0.74mg m^{-3}), NO_2 ($28.5 \mu\text{g m}^{-3}$), NO_x ($39.0 \mu\text{g m}^{-3}$), O_3 ($67.7 \mu\text{g m}^{-3}$)
 675 and O_x ($96.1 \mu\text{g m}^{-3}$).

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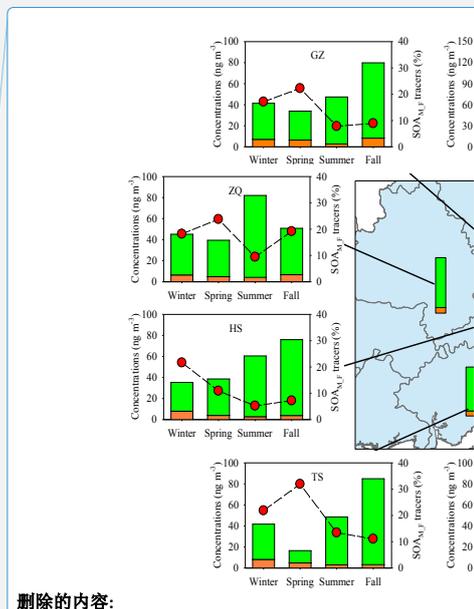




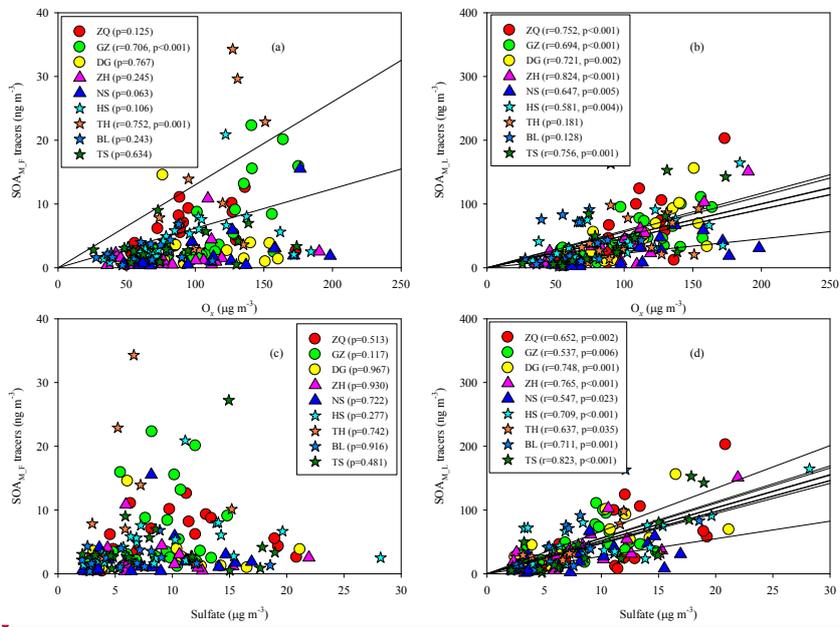
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679 Figure 4 Spatial and seasonal variations of SOA_M tracers at 9 sites in the PRD. The bars in the central figure represent
 680 the annual average concentrations of the SOA_M tracers.

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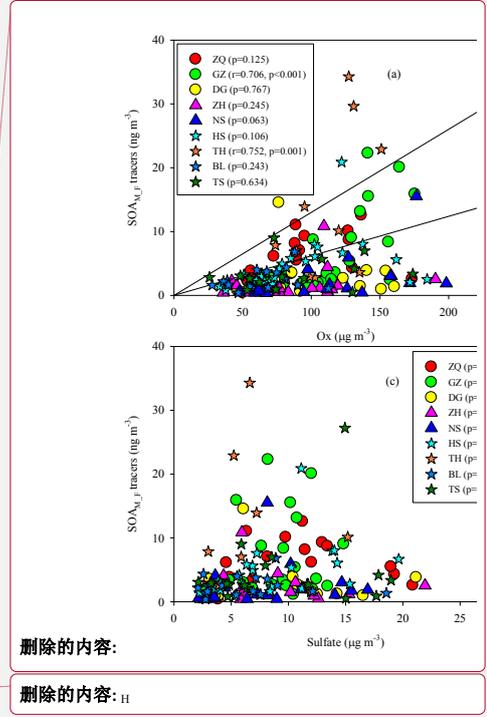
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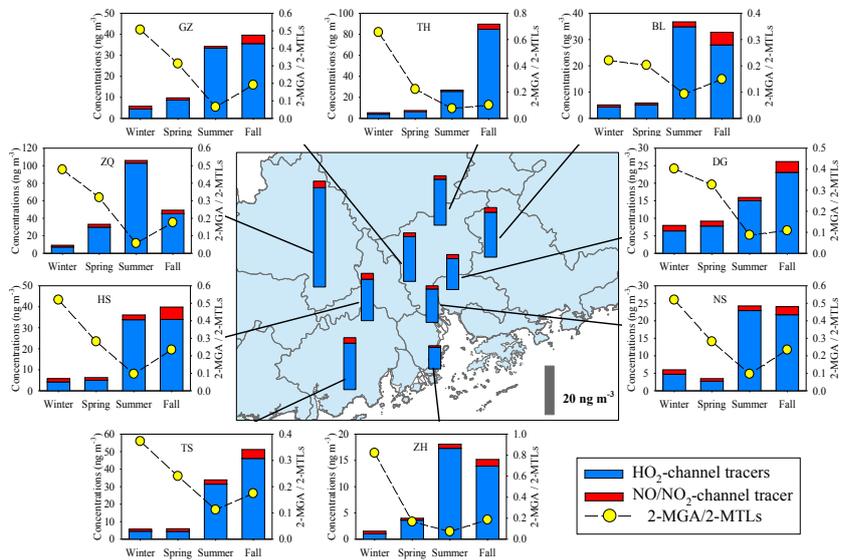


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684 Figure 5 Correlations of SOA_{M,F} and SOA_{M,L} tracers with O₃ (a, b) and sulfate (c, d)

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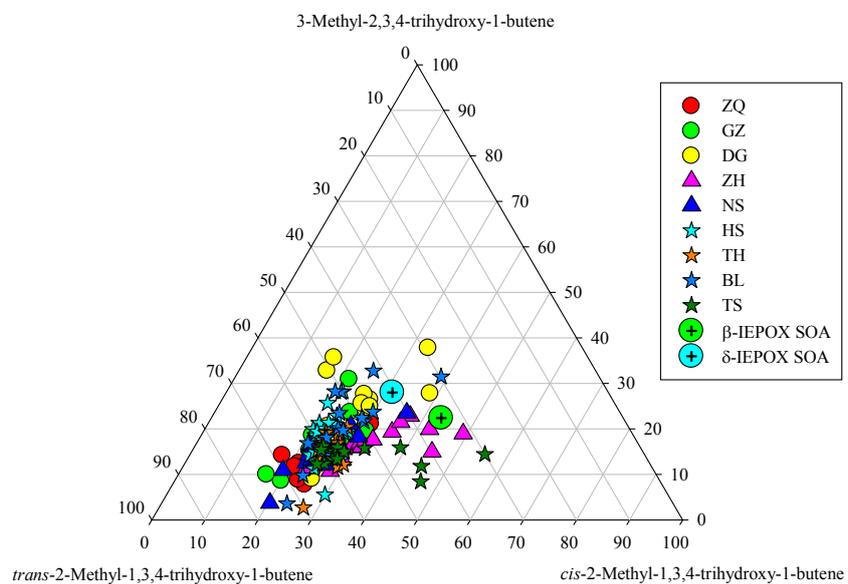




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689 Figure 6 Spatial and seasonal variations of SOA₁ tracers at 9 sites in the PRD. The bars in the central figure represent
 690 the annual average concentrations of the SOA₁ tracers.

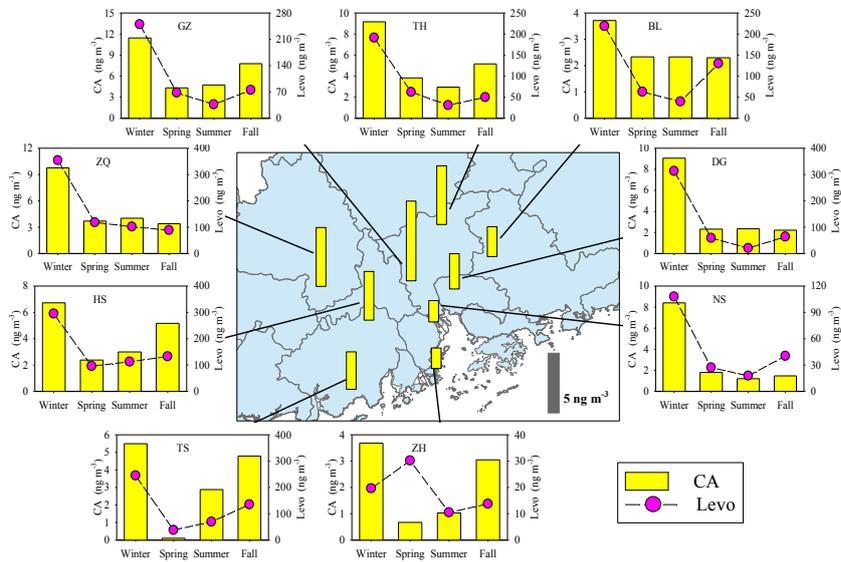
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693 Figure 7 Ternary plot of C₅-alkene triol isomers in the PRD samples and in the β-IEPOX and δ-IEPOX derived
 694 SOA (Lin et al., 2012).

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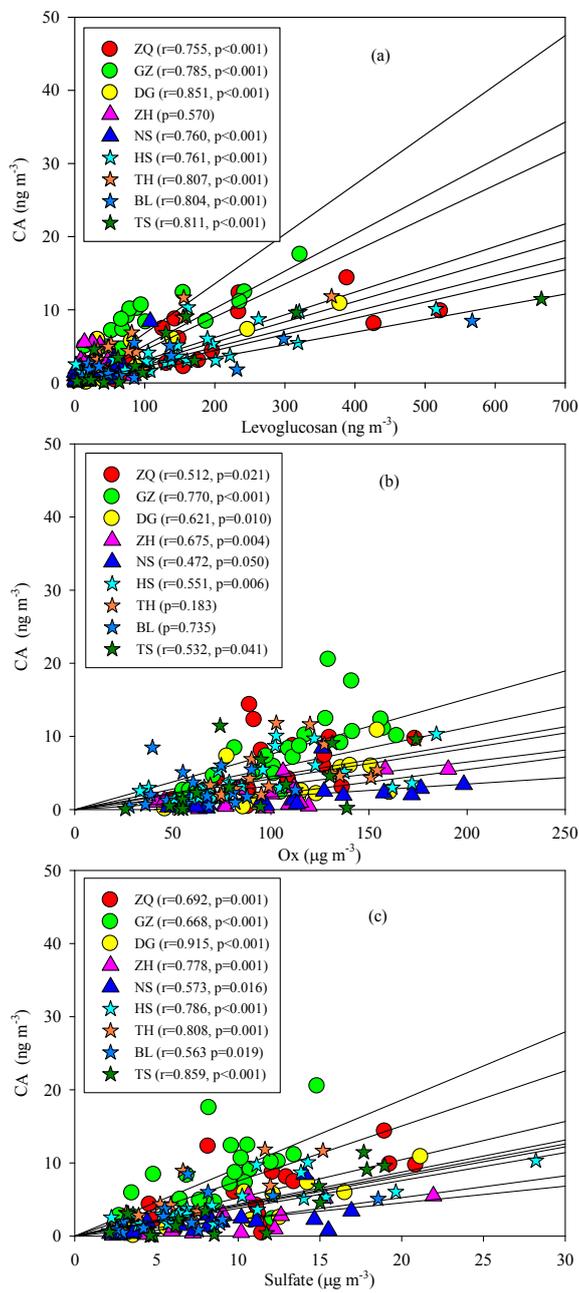
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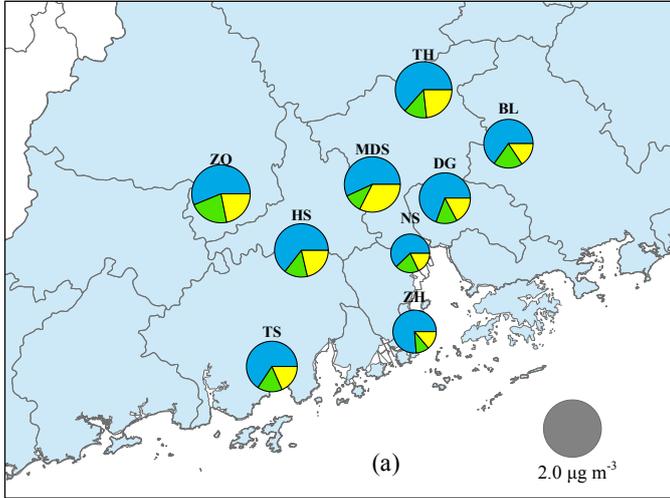
Figure 8 Spatial and seasonal variations of SOA_c tracer (CA) at 9 sites in the PRD. The bars in the central figure represent the annual average concentration of the SOA_c tracers. The pink circle indicates the BB tracer, levoglucosan (Levo).



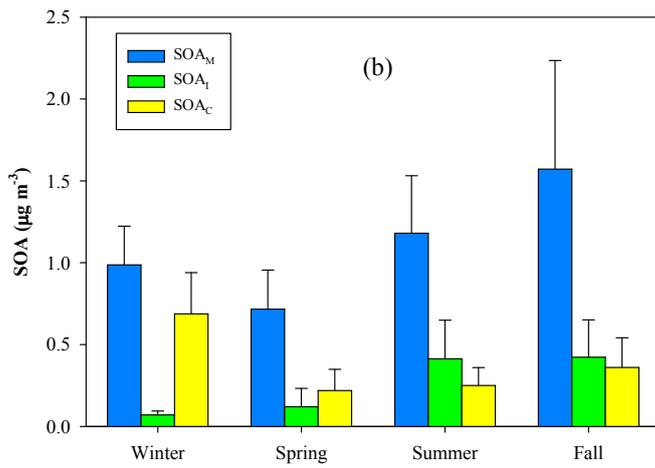
701

702 Figure 9 Significant correlations of CA with levoglucosan (a), Ox (b) and sulfate (c).

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707 Figure 10 Spatial (a) and seasonal variations (b) of BSOA

708

709 Table 1 Correlations between SOA_I tracers and NO_x

	2-MGA		2-MGA/2-MTLs	
	Coefficient (r)	<i>p</i> -value	Coefficient (r)	<i>p</i> -value
NO	0.028	0.733	0.166	0.043
NO ₂	0.205	0.008	0.352	<0.001
NO _x	0.132	0.102	0.286	<0.001
NO ₂ /NO	0.001	0.986	0.162	0.048

710

711

712 Table 2 Correlations of BSOA with sulfate and O_x

	Sulfate			O _x		
	Slope	<i>p</i> -value	% ^a	Slope	<i>p</i> -value	% ^a
SOA _M	0.112	<0.001	45	0.013	<0.001	57
SOA _I	0.020	<0.001	34	0.003	<0.001	50
SOA _C	0.041	<0.001	46	0.004	<0.001	55
BSOA	0.172	<0.001	44	0.019	<0.001	55

713 ^a Percentages of SOA reduction at 50% decline of sulfate or O_x

714