

We thank the reviewers for their time and insightful comments. Below are our responses to the comments. Reviewers' comments are italicized and the responses are underlined. When available, modified parts in the revised manuscript are presented in red. At the end, we have also included a list of minor modifications made in the revised manuscript.

Responses to Reviewer 1:

- 1. Line 221: The supposition that autooxidation is occurring deserves more support. For example, the statement concerning "a few hydroxy radical reaction reactions" should be expanded into a formal kinetic argument concerning the low probability of multiple OH-initiated oxidation steps.*

In the revised manuscript, we compared the characteristic time of OH oxidation reaction derived from the lifetime of α -pinene with the peak time of $C_{10}H_{17}NO_{\geq 6}$ in FIGAERO-CIMS, which supports that these compounds are likely formed via one OH oxidation reaction. With one OH oxidation reaction, $C_{10}H_{17}NO_{\geq 6}$ is likely an autoxidation product (Xu et al., 2019; Berndt et al., 2016; Pye et al., 2019).

"This observation is consistent with the autoxidation mechanism, in which highly oxidized molecules are formed in a short time scale (Ehn et al., 2014; Crouse et al., 2013; Jokinen et al., 2015). **Based on FIGAERO-HR-ToF-I-CIMS data (Fig. S1), $C_{10}H_{17}NO_{\geq 6}$ peak at 75 min. This is comparable to the lifetime of α -pinene at 53 min in the same experiment, which is related to characteristic time of OH oxidation. This suggests that the aforementioned ON are likely formed via one OH oxidation reaction, which is consistent with the autoxidation scheme to generate $C_{10}H_{17}NO_{\geq 6}$ proposed in prior studies (Berndt et al., 2016; Xu et al., 2019; Pye et al., 2019). It is important to note that...**"

- 2. Line 228: Again, the actual kinetic argument should be explicitly given. I assume that, combined with a typical $RO_2 + NO$ rate constant and 10 ppb NO, the predicted $RO_2 + NO$ pseudo first order rate constant is a "a few per second."*

We have included the typical rate constant of RO_2+NO in the revised manuscript.

"... that the autoxidation rate constant for the α -pinene+OH \cdot system could be up to a few per second, which is comparable to the NO level of ~ 10 ppb, **assuming a typical RO_2+NO reaction rate constant of $1 \times 10^{-11} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ (Orlando and Tyndall, 2012).**"

- 3. Line 348: Although this paragraph is making an appropriate detailed argument about the difference between the two proxies, I think it is important to point out very simply to the reader that the reason why the proxy $NO_3\text{-(org)}/Org$ works better is because the common atmospheric chamber problem of wall loss effectively gets cancelled out with this proxy method.*

We have expanded the discussion in the revised manuscript. Precisely, the argument that the use of proxy $NO_{3,Org}/Org$ cancels out the vapor wall loss effect may not be entirely valid because organic vapors get lost to the Teflon surface at a varying rate as a function of saturation mass concentration (Zhang et al., 2015). For the sake of simplicity, we assume in this study that the vapor wall loss rate of organics and organic nitrates is the same. This assumption is likely reasonable based on the thermal desorption profiles of CHO and CHON at the peak SOA growth that we have added in the revised SI (Fig. S5).

“For example, Huang et al. (2018) observe that the decay of isoprene hydroxyl nitrate depends on the reactor humidity. While SO_4 is practically non-volatile in the experimental condition of this study, both ON and non-nitrated organics could have some fractions of semi-volatile species whose vapors are subject to wall loss. Assuming a uniform loss rate of ON and non-nitrated organic vapors to the reactor walls, the effect of vapor wall loss could be effectively cancelled out in $\text{NO}_{3,\text{Org}}/\text{Org}$, but not in $\text{NO}_{3,\text{Org}}/\text{SO}_4$. This assumption is likely reasonable because the vapor wall loss rate is a function of saturation mass concentration (Zhang et al., 2015) and the average saturation mass concentrations of bulk ON and non-nitrated organic aerosol are similar based on the thermal desorption profiles in FIGAERO-HR-ToF-I-CIMS (Fig. S5). Therefore, $\text{NO}_{3,\text{Org}}/\text{Org}$ is a better proxy to infer hydrolysis of $p\text{ON}$ than others.”

4. *Line 381: It is quite clear from previous work that iodide CIMS sensitivity varies greatly for moderately oxygenated organics. Therefore, it is also quite likely that these species are underestimated in the present work.*

We certainly believe that sensitivity varies greatly among moderately oxygenated organic species due to iodide CIMS selectivity. However, the focus in the concerned paragraph starting at line 369 in the original manuscript is β -pinene+ NO_3 · SOA that is comprised of less oxidized organic species instead of moderately oxygenated organics. Therefore, in the original manuscript at Lines 380-382, we included a sentence to describe the underestimating nature of our FIGAERO-HR-ToF-I-CIMS data.

5. *Line 385: This really supports the appropriateness of the “reconstructive” kinetics shown in Figure 3b and should be explicitly pointed out.*

We have made the suggested modification in the revised manuscript.

“As illustrated in Fig. 4, the time series of $p\text{ON}/\text{OA}$ stabilizes fairly quickly in most of the experiments, regardless of RH and/or the phase state of seed aerosol, supporting the appropriateness of the reconstructive approach shown in Fig. 3b.”

6. *Line 397: It should be added that bulk studies also found that these reactions did not require low pH to be fast.*

We have made the suggested modification in the revised manuscript.

“Particle acidity is found to enhance hydrolysis rate of α -pinene hydroxy nitrate (Rindelaub et al., 2016), though no clear difference is observed here between experiments with AS and SA+MS seed particles (i.e., Exp. 5 and 2). It is worth noting that aqueous AS particles are not neutral but slightly acidic due to partitioning of ammonium to the gas phase once the particles enter the chamber (Gao et al., 2004). In Gao et al. (2004), the reported pH of aqueous AS particles is 4.6 and we expect a similar pH in our study. In Rindelaub et al. (2016), the reported hydrolysis lifetime is short at 1.3 h at pH of 4.0. On the other hand, previous studies have shown that isoprene-derived hydroxyl nitrates do not require low pH to undergo fast hydrolysis (Darer et al., 2011; Hu et al., 2011). Thus, $p\text{ON}$ formed from...”

7. *Line 219 (Figure S1): I think there must be something wrong with the symbols (or legend) here, as I don't understand what is actually being plotted. For example, does any green series correspond to any H17 compound and not just O8H17 (green squares)?*

Yes, all the combinations of $C_{10}H_{13,15,17}NO_{5-8}$ are plotted in Fig. S1 for consistency. To clarify, we have modified the legend of Fig. S1. We have also modified the figure caption as follows.

“FIGAERO-HR-ToF-I-CIMS time-series data of select organic nitrate aerosol. All combinations of $C_{10}H_{13,15,17}NO_{5-8}$ with different oxygen and nitrogen numbers are shown here. Different symbols correspond to compounds with different oxygen numbers, while different colors correspond to compounds with different nitrogen numbers. Shown are the data from Exp. 3 (α -pinene+OH \cdot), though very similar trends are observed for β -pinene+OH \cdot from Exp. 6.”

8. *Line 236: Replace “subtraction” with “abstraction”*

We have made the suggested modification in the revised manuscript.

9. *Line 278 and beyond: Nitrate, sulfate, and ammonium ions must be indicated by their proper chemical formulas, NO₃⁻, SO₄²⁻, and NH₄⁺, respectively, rather than the shorthand NO₃, SO₄, and NH₄ used in the manuscript. This is potentially particularly confusing as neutral NO₃ radical reactions are important formation routes for organonitrates.*

It is a common practice to use NO₃, SO₄, and NH₄ to refer to aerosol nitrate, sulfate, and ammonium for AMS data (Jimenez et al., 2003; Allan et al., 2004; Canagaratna et al., 2007). This is because NO₃, SO₄, and NH₄ measured by AMS are not necessarily all from ions (i.e., NO₃⁻, SO₄²⁻, and NH₄⁺) but can also come from covalently-bonded species, such as R-ONO₂, R-OSO₃H, and R-NH₂. For this reason, the use of NO₃⁻, SO₄²⁻, and NH₄⁺ in our manuscript would not be appropriate, where we expect a substantial contribution of ON to NO₃ signal. In the original manuscript, we had explicitly written out “nitrate radical” (rather than formula) throughout the manuscript in order to avoid the confusion. To further clarify this, we have decided to replace “nitrate radical” with NO₃ \cdot (with a dot to denote a radical), such that readers do not get confused between NO₃ (aerosol nitrate) and NO₃ \cdot (nitrate radical). To be consistent, other radicals, such as hydroxyl radical written as OH \cdot , now carry \cdot in the revised manuscript.

10. *Line 298: This sentence is confusing. I suggest “ \cdot : : on hydrolysis on pON that partitions to the aerosol due to condensation rather than dissolution.”*

Line 314: This sentence is confusing. I suggest “ \cdot : : is a reaction in which liquid water is a reactant, it is : : :”

Line 349: Typo: replace “rector” with “reactor”

Line 355: Typo: replace “technique” with “techniques”

Line 360: Isn't this a mass concentration? This should be clarified.

Line 394: Add “a” between “in” and “prior”

Line 395: Add “here” between “observed” and “between”

Line 401: Add “method” after “solution”

Line 445: Replace “several” with “a few”

We have made the suggested modifications in the revised manuscript.

Responses to Reviewer 2:

1. Page 6, line 133: Please elaborate. What “negligible impact” on the results concerning hydrolysis was there?

Regarding the injection order, we do not observe clear difference in pON/OA time-series data between Exp. 8 and 10 (both red filled circles), in which only difference is the injection order of precursor VOC and oxidant (Fig. 4c). Regarding the branching ratio of $\text{RO}_2\cdot$, as shown in Figure 4c, no clear difference in pON/OA time-series data is observed between $\text{RO}_2\cdot+\text{NO}_3\cdot$ -dominant (filled circle) and $\text{RO}_2\cdot+\text{HO}_2\cdot$ -dominant cases (cross) under low RH, dry AS and high RH, wet AS. We interpret these results as neither injection order nor RO_2 branching ratio in the case of α -pinene+ $\text{NO}_3\cdot$ affects the susceptibility of pON against hydrolysis.

2. Page 10, line 233: This is likely the case for beta-pinene + NO_3 , but alpha-pinene+ NO_3 leads to very low ON yields (preferably forming volatile pinonaldehyde) and low SOA yields as discussed in Fry et al. (2014) as well as for -terpinene + NO_3 as discussed in Slade et al. (2017). Favorability of nitrated to non-nitrated organics following NO_3 oxidation vs. OH oxidation in the presence of NO will depend on the degree of branching on the carbon of the C-ONO2 bond. As stated, it is a bit oversimplified, please revise accordingly.

Concerning my second comment, it would be helpful if the authors included a discussion on the difference in the degree of nitration between the beta-pinene and alphapinene systems under the “nighttime” scenario when oxidized by NO_3 . The current discussion focuses more on differences between OH and NO_3 oxidation and very little on differences between the precursor VOCs. The authors do say on page 9, lines 195-197 that a more distinct difference is observed between different oxidation conditions than between different precursor VOCs, but I would expect, based on previous literature, that substantial differences in CHO and CHNO fractions would exist between alpha and beta-pinene owing to the positions of their double bonds. Please elaborate.

Indeed, the total organic nitrate yield in $\text{NO}_3\cdot$ oxidation of α -pinene and γ -terpinene is low, compared to $\text{NO}_3\cdot$ oxidation of other monoterpenes. However, the corresponding section (i.e., Section 3.1) is focused on the chemical composition of particles rather than the total products in both gas and particle phases. Nevertheless, we have modified the corresponding paragraphs to clarify that our observation of a high abundance of ON is not in a direct contradiction to previously reported low ON yield, if low-volatility species contributing to SOA is composed of ON.

“This observation is consistent with a direct addition of a nitrate functional group to a double bond (Wayne et al., 1991; Ng et al., 2017), whereas the formation of ON in OH \cdot oxidation condition is a minor channel of $\text{RO}_2\cdot+\text{NO}$ reaction (Perring et al., 2013). Although this is generally true for many monoterpenes, the organic nitrate yield in α -pinene+ $\text{NO}_3\cdot$ has been known to be low (Fry et al., 2014) due to loss of a nitrate functional group followed by alkoxy radical bond scission (Kurten et al., 2017). Fry et al. (2014) observe no SOA formation in the same system, though another study has reported the formation of non-negligible SOA mass even at a relatively low initial concentration of α -pinene (Nah et al., 2016). Since the particle-phase compounds represent <10% of overall α -pinene+ $\text{NO}_3\cdot$ products by mass, it is not necessarily inconsistent to observe a higher abundance of ON than non-nitrated organics in the particle phase if low-volatility compounds mainly consist of ON. Moreover, the contribution...”

3. Page 10, line 236: *I don't quite understand this statement. Both OH and NO₃ can abstract hydrogens, but ultimately it depends on the degree of saturation, or? OH and NO₃ preferably add to carbon-carbon double bonds over hydrogen abstraction. Please rephrase.*

We have made the suggested modification in the revised manuscript.

“Moreover, the contribution from species containing C_{<10} is minimal in the NO₃· oxidation condition. Once NO₃· attacks a double bond in the initial oxidation reaction of α-pinene and β-pinene, the reaction products no longer contain any double bond. Unlike OH·, hydrogen abstraction reaction by NO₃· is slower by orders of magnitude (Atkinson and Arey, 2003). Therefore, multi-generation oxidation is unlikely to occur within the timescale of experiments (Wayne et al., 1991; Ng et al., 2017). This means that once the precursor VOC undergoes functionalization upon the initial NO₃· oxidation, it is not likely to experience fragmentation during the experiment.”

4. *There is no mention of pH-dependent (i.e., acid-catalyzed) hydrolysis and the formation of organic sulfates (Liggio and Li, 2008; Rindelaub et al., 2016; Rindelaub et al., 2015). This is likely an important process affecting pON concentrations and the NO₃,Org/SO₄ ratio. Could you estimate what fraction of SO₄ is organic?*

Regarding pH-dependent hydrolysis rate, in the original manuscript, we had briefly discussed previous studies on dependence of the hydrolysis rate on acidity (lines 394-397 of original manuscript). In response to comment #6 from Reviewer 1, we have made the suggested modification in the revised manuscript.

“Particle acidity is found to enhance hydrolysis rate of α-pinene hydroxy nitrate (Rindelaub et al., 2016), though no clear difference is observed here between experiments with AS and SA+MS seed particles (i.e., Exp. 5 and 2). It is worth noting that aqueous AS particles are not neutral but slightly acidic due to partitioning of ammonium to the gas phase once the particles enter the chamber (Gao et al., 2004). In Gao et al. (2014), the reported pH of aqueous AS particles is 4.6 and we expect a similar pH in our study. In Rindelaub et al. (2016), the reported hydrolysis lifetime is short at 1.3 h at pH of 4.0. On the other hand, previous studies have shown that isoprene-derived hydroxyl nitrates do not require low pH to undergo fast hydrolysis (Darer et al., 2011; Hu et al., 2011). Thus, pON formed from...”

Indeed, the formation of organic sulfate has been reported in past studies. However, the extent to which organic sulfate is formed will not affect our interpretation of the time-series data of NO₃,Org/SO₄ ratio (Fig. 3). In AMS, the majority of organic sulfate (>96%) is detected as organic (C_xH_yO_z⁺) and sulfate (H_xSO_y⁺) fragments rather than organosulfur fragments (C_xH_yO_zS⁺) (Chen et al., 2019). Thus, the conversion of inorganic to organic sulfates hardly affects the absolute value of SO₄ measured by AMS.

5. Page 14, line 330: *The difference between the changes in NO₃,Org/SO₄ due to aerosol water content and reactor humidity alone is not clear. Doesn't the reactor humidity affect aerosol water content and subsequently pON hydrolysis? Please clarify how "reactor humidity alone" and not water content affects NO₃,Org/SO₄.*

Previous chamber studies (Liu et al., 2012; Boyd et al., 2015; Rindelaub et al., 2016) have investigated hydrolysis of pON by comparing results from low RH experiment (<10%) with those from high RH experiment (>40%). The issue of this approach is that chamber RH does not only affect aerosol water content but also absolute water vapor pressures (gas-phase chemistry) as well as other factors impacted by

RH, such as loss rates of organic vapors to chamber walls, as discussed in the response to comment #3 from Reviewer 1. In this study, we performed an experiment with high reactor RH but with effloresced seed aerosols to isolate the effects stemming from differences in water vapor pressures or vapor wall loss, to which we refer as “reactor humidity alone.” As shown in Fig. 3, it is evident that the decay behavior of $\text{NO}_{3,\text{Org}}/\text{SO}_4$ from high RH experiments with effloresced (Exp. 4) and deliquesced seed aerosols (Exp. 5) is very similar, suggesting that changes in the time evolution of $\text{NO}_{3,\text{Org}}/\text{SO}_4$ between low and high RH experiments do not stem from changes in the aerosol water content but rather from changes in other factors impacted by RH, such as vapor wall loss.

6. *Figure 3b: The y-axis label is confusing “Exp.4/Exp.3”, are you dividing results from Exp.4 by Exp.3? Please modify.*

Yes, each proxy in Exp. 4 is divided by that in Exp. 3 to express a relative decay between high and low RH experiments. We have modified the y-axis label of Fig. 3b and have added a sentence in the caption to clarify the y-axis.

“Figure 3. (a) Time-series data of $\text{NO}_{3,\text{Org}}/\text{SO}_4$ from Exp. 3-5 (α -pinene+OH \cdot) and the exponential fits with corresponding characteristic times. (b) $\text{NO}_{3,\text{Org}}/\text{Org}$, $\text{NO}_{3,\text{Org}}/\text{SO}_4$, reconstructed $\text{NO}_{3,\text{Org}}/\text{Org}$ based on the decay rate of $\text{NO}_{3,\text{Org}}/\text{SO}_4$. **Each proxy in Exp. 4 is divided by that in Exp. 3 to determine the relative decay between high and low RH experiments.**”

7. *Page 14, discussion on volatility: another interpretation of the $\text{NO}_3,\text{Org}/\text{Org}$ and $\text{NO}_3,\text{Org}/\text{SO}_4$ decay curves is the organosulfate products that possibly form are sufficiently low volatility that they remain in the aerosol phase, but a fraction of the hydrolyzed ON that is not organosulfate is sufficiently volatile that it repartitions to the gas phase.*

We have included a sentence indicating the possibility of the formation of organic sulfate as a way to generate sufficiently low volatile organics in the revised manuscript.

“This is because the organic moiety of hydrolysis product is generally considered to have a substituted alcohol group (Boschan et al., 1955) and to have a relatively similar vapor pressure to the ON compound and shall remain in the particle phase (Pankow and Asher, 2008). **Similar to this is the formation of organic sulfate from hydrolysis of pON (Liggio and Li, 2006, 2008; Surratt et al., 2008), which has sufficiently low volatility to remain in the particle phase as with alcohol-substituted products.** However, the measured...”

8. *Page 17, lines 394-395: If no difference in pON hydrolysis is observed between AS and SA+MS seed experiments, might it be that HNO_3 in the particles sufficiently decreases aerosol pH that both “neutral” AS seed experiments and “acidic” SA+MS seed experiments are essentially the same pH? Aerosol pH measurements, e.g., with pH paper as in Craig et al. (2018), or predictions via AIM or ISORROPIA would help with interpretation but may be beyond the scope of this study.*

This is an interesting perspective on the effect of partitioned HNO_3 . First, aqueous AS seed particles are not “neutral” but is already acidic due to partitioning of ammonium to the gas phase once the particles entered the chamber (Gao et al., 2004), as mentioned in the response to comment #6 from Reviewer 1. Second, while HNO_3 could potentially lower aerosol pH down to ~1-2 (Guo et al., 2016), it is not likely

reach the same level of pH as SA+MS seed particles (Gao et al., 2004). A more comprehensive study on pH-dependence on the hydrolysis rate as well as hydrolyzable fraction of organic nitrate under different pH conditions will be required.

9. *Page 17, line 415: Please include citation. Presumably, the authors refer to greater uptake of N₂O₅ at higher relative humidity due to changes in particle viscosity? A good example would be Grzanic et al. (2015).*

We have included the suggested citation in the revised manuscript.

“The reduction of NO₃/Org in Fig. S5 (~30 %) is greater than in Boyd et al. (2015) (10 %), which could be because (1) the amount of N₂O₅, a source of inorganic nitrate, used in our study is slightly larger and (2) RH in our study is higher by 10-20 %, which may have allowed greater uptake of N₂O₅ to produce inorganic nitrate **due to changes in aerosol viscosity (Grzanic et al., 2015).**”

Additional modifications:

- Typo correction
 - Line 181: Ccondensation → **Condensation**
 - Line 213: are → **is**
 - Line 437: pahse → **phase**
 - Line 60 in SI: $C_{<9} \text{pON}$ → $C_{\leq 9} \text{pON}$
- Figures and tables
 - RH reflecting the exact measurement values has been updated in Table 1.
 - A sentence has been added in the caption of Fig. 5 to clarify the use of different font sizes in the figure.
 - Label and unit on y-axis in Fig. 6 have been added.
- Clarification
 - The sentence describing the use of the unity collection efficiency for the AMS data has been added.
 - The sentence describing the use of uniform sensitivity among detected species in FIGAERO-HR-ToF-I-CIMS has been added.
 - Eq. 1 has been updated to include another form to express pON/OA.

References

Allan, J. D., Delia, A. E., Coe, H., Bower, K. N., Alfarra, M. R., Jimenez, J. L., Middlebrook, A. M., Drewnick, F., Onasch, T. B., Canagaratna, M. R., Jayne, J. T., and Worsnop, D. R.: A generalised method for the extraction of chemically resolved mass spectra from aerodyne aerosol mass spectrometer data, *J Aerosol Sci*, 35, 909-922, 10.1016/j.jaerosci.2004.02.007, 2004.

Atkinson, R., and Arey, J.: Atmospheric degradation of volatile organic compounds, *Chem Rev*, 103, 4605-4638, 10.1021/cr0206420, 2003.

Berndt, T., Richters, S., Jokinen, T., Hyttinen, N., Kurten, T., Otkjaer, R. V., Kjaergaard, H. G., Stratmann, F., Herrmann, H., Sipila, M., Kulmala, M., and Ehn, M.: Hydroxyl radical-induced formation of highly oxidized organic compounds, *Nat Commun*, 7, 10.1038/ncomms13677, 2016.

Boschan, R., Merrow, R. T., and Vandolah, R. W.: The Chemistry of Nitrate Esters, *Chem Rev*, 55, 485-510, DOI 10.1021/cr50003a001, 1955.

Boyd, C. M., Sanchez, J., Xu, L., Eugene, A. J., Nah, T., Tuet, W. Y., Guzman, M. I., and Ng, N. L.: Secondary organic aerosol formation from the beta-pinene+NO₃ system: effect of humidity and peroxy radical fate, *Atmos Chem Phys*, 15, 7497-7522, 10.5194/acp-15-7497-2015, 2015.

Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q., Onasch, T. B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R., Trimborn, A. M., Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits, P., and Worsnop, D. R.: Chemical and microphysical characterization of ambient aerosols with the aerodyne aerosol mass spectrometer, *Mass Spectrom Rev*, 26, 185-222, 10.1002/mas.20115, 2007.

Chen, Y. L., Xu, L., Humphry, T., Hettiyadura, A. P. S., Ovadnevaite, J., Huang, S., Poulain, L., Schroder, J. C., Campuzano-Jost, P., Jimenez, J. L., Herrmann, H., O'Dowd, C., Stone, E. A., and Ng, N. L.: Response of the Aerodyne Aerosol Mass Spectrometer to Inorganic Sulfates and Organosulfur Compounds: Applications in Field and Laboratory Measurements, *Environ Sci Technol*, 53, 5176-5186, 10.1021/acs.est.9b00884, 2019.

Crouse, J. D., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., and Wennberg, P. O.: Autoxidation of Organic Compounds in the Atmosphere, *J Phys Chem Lett*, 4, 3513-3520, 10.1021/jz4019207, 2013.

Ehn, M., Thornton, J. A., Kleist, E., Sipila, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I. H., Rissanen, M., Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurten, T., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., Canagaratna, M., Dal Maso, M., Berndt, T., Petaja, T., Wahner, A., Kerminen, V. M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol, *Nature*, 506, 476+, 10.1038/nature13032, 2014.

Fry, J. L., Draper, D. C., Barsanti, K. C., Smith, J. N., Ortega, J., Winkle, P. M., Lawler, M. J., Brown, S. S., Edwards, P. M., Cohen, R. C., and Lee, L.: Secondary Organic Aerosol Formation and Organic Nitrate Yield from NO₃ Oxidation of Biogenic Hydrocarbons, *Environ Sci Technol*, 48, 11944-11953, 10.1021/es502204x, 2014.

Gao, S., Ng, N. L., Keywood, M., Varutbangkul, V., Bahreini, R., Nenes, A., He, J. W., Yoo, K. Y., Beauchamp, J. L., Hodyss, R. P., Flagan, R. C., and Seinfeld, J. H.: Particle phase acidity and oligomer formation in secondary organic aerosol, *Environ Sci Technol*, 38, 6582-6589, 10.1021/es049125k, 2004.

Grzanic, G., Bartels-Rausch, T., Berkemeier, T., Turler, A., and Ammann, M.: Viscosity controls humidity dependence of N₂O₅ uptake to citric acid aerosol, *Atmos Chem Phys*, 15, 13615-13625, 10.5194/acp-15-13615-2015, 2015.

Guo, H., Sullivan, A. P., Campuzano-Jost, P., Schroder, J. C., Lopez-Hilfiker, F. D., Dibb, J. E., Jimenez, J. L., Thornton, J. A., Brown, S. S., Nenes, A., and Weber, R. J.: Fine particle pH and the partitioning of nitric acid during winter in the northeastern United States, *J Geophys Res-Atmos*, 121, 10355-10376, 10.1002/2016jd025311, 2016.

Jimenez, J. L., Jayne, J. T., Shi, Q., Kolb, C. E., Worsnop, D. R., Yourshaw, I., Seinfeld, J. H., Flagan, R. C., Zhang, X. F., Smith, K. A., Morris, J. W., and Davidovits, P.: Ambient aerosol sampling using the Aerodyne Aerosol Mass Spectrometer, *J Geophys Res-Atmos*, 108, Artn 8425, 10.1029/2001jd001213, 2003.

Jokinen, T., Berndt, T., Makkonen, R., Kerminen, V. M., Junninen, H., Paasonen, P., Stratmann, F., Herrmann, H., Guenther, A. B., Worsnop, D. R., Kulmala, M., Ehn, M., and Sipila, M.: Production of extremely low volatile organic compounds from biogenic emissions: Measured yields and atmospheric implications, *P Natl Acad Sci USA*, 112, 7123-7128, 10.1073/pnas.1423977112, 2015.

Kurten, T., Moller, K. H., Nguyen, T. B., Schwantes, R. H., Misztal, P. K., Su, L. P., Wennberg, P. O., Fry, J. L., and Kjaergaard, H. G.: Alkoxy Radical Bond Scissions Explain the Anomalously Low Secondary Organic Aerosol and Organonitrate Yields From alpha-Pinene + NO₃, *J Phys Chem Lett*, 8, 2826-2834, 10.1021/acs.jpcllett.7b01038, 2017.

Liggio, J., and Li, S. M.: Reactive uptake of pinonaldehyde on acidic aerosols, *J Geophys Res-Atmos*, 111, Artn D24303, 10.1029/2005jd006978, 2006.

Liggio, J., and Li, S. M.: Reversible and irreversible processing of biogenic olefins on acidic aerosols, *Atmos Chem Phys*, 8, 2039-2055, DOI 10.5194/acp-8-2039-2008, 2008.

Liu, S., Shilling, J. E., Song, C., Hiranuma, N., Zaveri, R. A., and Russell, L. M.: Hydrolysis of Organonitrate Functional Groups in Aerosol Particles, *Aerosol Sci Tech*, 46, 1359-1369, 10.1080/02786826.2012.716175, 2012.

Nah, T., Sanchez, J., Boyd, C. M., and Ng, N. L.: Photochemical Aging of alpha-pinene and beta-pinene Secondary Organic Aerosol formed from Nitrate Radical Oxidation, *Environ Sci Technol*, 50, 222-231, 10.1021/acs.est.5b04594, 2016.

Ng, N. L., Brown, S. S., Archibald, A. T., Atlas, E., Cohen, R. C., Crowley, J. N., Day, D. A., Donahue, N. M., Fry, J. L., Fuchs, H., Griffin, R. J., Guzman, M. I., Herrmann, H., Hodzic, A., Iinuma, Y., Jimenez, J. L., Kiendler-Scharr, A., Lee, B. H., Luecken, D. J., Mao, J. Q., McLaren, R., Mutzel, A., Osthoff, H. D., Ouyang, B., Picquet-Varrault, B., Platt, U., Pye, H. O. T., Rudich, Y., Schwantes, R. H., Shiraiwa, M., Stutz, J., Thornton, J. A., Tilgner, A., Williams, B. J., and Zaveri, R. A.: Nitrate radicals and biogenic volatile

organic compounds: oxidation, mechanisms, and organic aerosol, *Atmos Chem Phys*, 17, 2103-2162, 10.5194/acp-17-2103-2017, 2017.

Pankow, J. F., and Asher, W. E.: SIMPOL.1: a simple group contribution method for predicting vapor pressures and enthalpies of vaporization of multifunctional organic compounds, *Atmos Chem Phys*, 8, 2773-2796, DOI 10.5194/acp-8-2773-2008, 2008.

Perring, A. E., Pusede, S. E., and Cohen, R. C.: An Observational Perspective on the Atmospheric Impacts of Alkyl and Multifunctional Nitrates on Ozone and Secondary Organic Aerosol, *Chem Rev*, 113, 5848-5870, 10.1021/cr300520x, 2013.

Pye, H. O. T., D'Ambro, E. L., Lee, B., Schobesberger, S., Takeuchi, M., Zhao, Y., Lopez-Hilfiker, F., Liu, J. M., Shilling, J. E., Xing, J., Mathur, R., Middlebrook, A. M., Liao, J., Welti, A., Graus, M., Warneke, C., de Gouw, J. A., Holloway, J. S., Ryerson, T. B., Pollack, I. B., and Thornton, J. A.: Anthropogenic enhancements to production of highly oxygenated molecules from autoxidation, *P Natl Acad Sci USA*, 116, 6641-6646, 10.1073/pnas.1810774116, 2019.

Rindelaub, J. D., Borca, C. H., Hostetler, M. A., Slade, J. H., Lipton, M. A., Slipchenko, L. V., and Shepson, P. B.: The acid-catalyzed hydrolysis of an alpha-pinene-derived organic nitrate: kinetics, products, reaction mechanisms, and atmospheric impact, *Atmos Chem Phys*, 16, 15425-15432, 10.5194/acp-16-15425-2016, 2016.

Surratt, J. D., Gomez-Gonzalez, Y., Chan, A. W. H., Vermeulen, R., Shahgholi, M., Kleindienst, T. E., Edney, E. O., Offenberg, J. H., Lewandowski, M., Jaoui, M., Maenhaut, W., Claeys, M., Flagan, R. C., and Seinfeld, J. H.: Organosulfate formation in biogenic secondary organic aerosol, *J Phys Chem A*, 112, 8345-8378, 10.1021/jp802310p, 2008.

Wayne, R. P., Barnes, I., Biggs, P., Burrows, J. P., Canosamas, C. E., Hjorth, J., Lebras, G., Moortgat, G. K., Perner, D., Poulet, G., Restelli, G., and Sidebottom, H.: The Nitrate Radical - Physics, Chemistry, and the Atmosphere, *Atmos Environ a-Gen*, 25, 1-203, Doi 10.1016/0960-1686(91)90192-A, 1991.

Xu, L., Moller, K. H., Crouse, J. D., Otkjwr, R. V., Kjaergaard, H. G., and Wennberg, P. O.: Unimolecular Reactions of Peroxy Radicals Formed in the Oxidation of alpha-Pinene and beta-Pinene by Hydroxyl Radicals, *J Phys Chem A*, 123, 1661-1674, 10.1021/acs.jpca.8b11726, 2019.

Zhang, X., Schwantes, R. H., McVay, R. C., Lignell, H., Coggon, M. M., Flagan, R. C., and Seinfeld, J. H.: Vapor wall deposition in Teflon chambers, *Atmos Chem Phys*, 15, 4197-4214, 10.5194/acp-15-4197-2015, 2015.

1 **Chemical Composition and Hydrolysis of Organic Nitrate Aerosol Formed from Hydroxyl and**
2 **Nitrate Radical Oxidation of α -pinene and β -pinene**

3

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13 **Keywords**

14 Aerosol, chemical composition, hydrolysis, monoterpenes, organic nitrate, hydroxyl and nitrate radical
15 oxidation

16

17 **Abstract**

18 Atmospheric organic nitrate (ON) is thought to play a crucial role in the formation potential of
19 ozone and aerosol, which are the leading air pollutants of concern across the world. Limited fundamental
20 knowledge and understanding of the life cycles of ON currently hinders the ability to quantitatively assess
21 its impacts on the formation of these pollutants. Although hydrolysis is currently considered as an important
22 loss mechanism of ON based on prior field measurement studies, this process for atmospherically relevant
23 ON has not been well constrained by fundamental laboratory studies. In this comprehensive study, we
24 investigated the chemical composition and hydrolysis process of particulate ON (p ON) formed from the
25 oxidation of α -pinene and β -pinene by hydroxyl ($\text{OH}\cdot$) and nitrate radicals ($\text{NO}_3\cdot$). For p ON that undergoes
26 hydrolysis, the hydrolysis lifetime is determined to be no more than 30 min for all systems explored. This
27 is significantly shorter than those reported in previous chamber studies (i.e., 3-6 h) but is consistent with
28 the reported lifetime from bulk solution measurement studies (i.e., 0.02-8.8 h). The discrepancy appears to
29 stem from the choice of proxy used to estimate the hydrolysis lifetime. The measured hydrolyzable fractions
30 of p ON (F_H) in the α -pinene+ $\text{OH}\cdot$, β -pinene+ $\text{OH}\cdot$, α -pinene+ $\text{NO}_3\cdot$, and β -pinene+ $\text{NO}_3\cdot$ systems are 23-32,
31 27-34, 9-17, and 9-15 %, respectively. While a very low F_H for the $\text{NO}_3\cdot$ ~~nitrate radical~~ oxidation system is
32 expected based on prior studies, F_H for the ~~hydroxyl radical~~ $\text{OH}\cdot$ oxidation system is surprisingly lower than
33 predicted in past studies. Overall, the hydrolysis lifetime as well as F_H obtained in this study serve as
34 experimentally constrained parameters that are required in regional and global chemical transport models
35 to accurately evaluate the impacts of ON on nitrogen budget and formation of ozone and aerosol.

36

37 1. Introduction

38 The oxidation of biogenic volatile organic compounds (BVOC) by ozone (O_3), hydroxyl radicals
39 ($OH\cdot$) and nitrate radicals ($NO_3\cdot$) is a major source of secondary organic aerosol (SOA) globally
40 (Kanakidou et al., 2005; Goldstein and Galbally, 2007; Spracklen et al., 2011). Many studies have pointed
41 to the synergistic effects of anthropogenic emissions on biogenic SOA formation in the atmosphere (Weber
42 et al., 2007; Carlton et al., 2010; Hoyle et al., 2011; Shilling et al., 2013; Xu et al., 2015a; Shrivastava et
43 al., 2019). The oxidation of BVOC in environments with anthropogenic NO_x emissions is an important
44 mechanism for coupled biogenic-anthropogenic interactions. In the presence of NO_x , the oxidation of
45 BVOC can lead to the formation of organic nitrate (ON), a large component of reactive oxidized nitrogen.
46 Results from ambient field measurements have revealed the ubiquitous presence of particulate ON (pON),
47 where it contributes to a large fraction of submicron organic aerosol at different sites worldwide (Fry et al.,
48 2013; Xu et al., 2015b; Liu et al., 2012a; Rollins et al., 2012; Rollins et al., 2013; Lee et al., 2016; Kiendler-
49 Scharr et al., 2016; Ng et al., 2017). These findings highlights the importance to understand the formation
50 and fates of ON to accurately evaluate its roles in NO_x recycling, O_3 , and SOA formation.

51 Monoterpenes ($C_{10}H_{16}$) is a major class of BVOC, with annual emissions of 157-177 Tg C yr^{-1}
52 (Guenther et al., 2012). Laboratory studies have demonstrated that the ~~$NO_3\cdot$ nitrate radical~~ oxidation of
53 monoterpenes leads to a substantial formation of ON and SOA, with ON yields up to ~70% (Wangberg et
54 al., 1997; Berndt and Boge, 1997; Griffin et al., 1999; Hallquist et al., 1999; Spittler et al., 2006; Fry et al.,
55 2009; Fry et al., 2014; Boyd et al., 2015; Nah et al., 2016; Boyd et al., 2017; Slade et al., 2017; Clafin and
56 Ziemann, 2018). For photooxidation of monoterpenes in the presence of NO_x , ON yields as high as 26%
57 have been reported for α -pinene (Noziere et al., 1999; Aschmann et al., 2002; Rindelaub et al., 2015).
58 Monoterpene emissions do not depend strongly on light and typically continue at night, making them
59 important ON and SOA precursors at any times of the day (daytime and nighttime) and throughout the year
60 (different seasons). It has been shown that monoterpene-derived ON is prevalent in areas where there are
61 substantial biogenic-anthropogenic interactions and oxidation of monoterpenes contributes to a large

62 fraction of SOA observed in the Southeastern U.S. (Xu et al., 2015a; Xu et al., 2015b; Lee et al., 2016;
63 Zhang et al., 2018; Xu et al., 2018a).

64 One of the largest uncertainties in our understanding of monoterpene ON chemistry is the extent to
65 which ON act as a permanent sink versus temporary reservoir of NO_x (Takeuchi and Ng, 2018). This would
66 depend on the fates of ON as they can either retain or release NO_x upon further reactions. Once formed,
67 gas-phase ON can undergo photolysis or ~~hydroxyl radical~~OH· oxidation to release NO_x or partition into the
68 particle phase. _pON in turn can undergo further chemistry to release NO_x or hydrolyze in the particle phase
69 to form nitric acid (HNO₃). Further, ON and HNO₃ can be removed via dry and wet deposition. One
70 important reaction of ON in the particle phase is hydrolysis in the presence of aerosol water, which is a
71 mechanism for NO_x loss (Day et al., 2010; Russell et al., 2011). Studies with bulk solutions showed that
72 particle-phase hydrolysis of tertiary nitrate is fast with a lifetime of minutes, while primary and secondary
73 nitrate is stable (Darer et al., 2011; Hu et al., 2011). However, the hydrolysis of _pON in aerosol water is
74 largely unconstrained. Results from field and modeling studies suggested a _pON lifetime of a few hours
75 (Pye et al., 2015; Lee et al., 2016; Fisher et al., 2016; Zare et al., 2018). A few recent laboratory chamber
76 studies elicited a complex picture where _pON formed from photochemical oxidation and ~~NO₃· nitrate radical~~
77 oxidation of monoterpenes appear to experience different magnitudes of hydrolysis (Boyd et al., 2015;
78 Rindelaub et al., 2015; Bean and Hildebrandt Ruiz, 2016; Boyd et al., 2017), likely due to the difference in
79 the relative amount of primary, secondary, and tertiary nitrate in these oxidation systems. Overall, there are
80 very limited studies on the further evolutions of ON produced from the oxidation of monoterpenes.

81 Here, we present results from a laboratory chamber study on the chemical composition and
82 hydrolysis process of _pON formed from oxidation of α-pinene and β-pinene by ~~OH· hydroxyl~~ and
83 ~~NO₃· nitrate radicals~~. Specifically, we report the hydrolysis lifetimes and the fraction of hydrolyzable _pON
84 formed in the systems examined in this study. This comprehensive chamber study on the hydrolysis of _pON
85 produced from various oxidation pathways of monoterpenes and peroxy radical (RO₂·) fates provides the
86 fundamental data to better constrain the role of hydrolysis in modulating _pON concentrations and lifetimes

87 in the atmosphere, their potential as a NO_x loss pathway, and their impacts on overall nitrogen budget, O₃
88 and SOA formation.

89

90 **2. Methods**

91 2.1. Chamber experiment design and procedure

92 A series of chamber experiments were performed in the Georgia Tech Environmental Chamber
93 facility (Boyd et al., 2015) housing two 12 m³ Teflon reactors. Precursor volatile organic compounds (VOC)
94 were α -pinene (99 %, Sigma-Aldrich) and β -pinene (99 %, Sigma-Aldrich) and the oxidation conditions of
95 interest were ~~OH·hydroxyl~~ and ~~NO₃·nitrate radical~~ oxidation, which were represented as “daytime” and
96 “nighttime” experiments, respectively. Specifically, four different systems of VOC and oxidation
97 conditions were studied: α -pinene+OH₂, β -pinene+OH₂, α -pinene+NO₃, and β -pinene+NO₃. In order to
98 infer the hydrolysis process, experiments were performed under low RH (i.e., ~5 %) or high RH (i.e., ~50-
99 70 %) conditions and with effloresced or deliquesced seed particles for the same initial concentrations of
100 precursor VOC and oxidant precursors. Temperature in the reactors was kept at room temperature (22-
101 25°C). Experimental conditions are summarized in Table 1.

102 Prior to every experiment, the reactor was flushed with zero air (AADCO, 747-14) for at least a
103 day. A typical experiment began with the injection of seed aerosol into the reactor by atomizing dilute
104 ammonium sulfate (AS; 0.015 M) or sulfuric acid + magnesium sulfate (SA+MS; 0.01 + 0.005 M) aqueous
105 solution. The seed aerosol was either directly atomized into the reactor or passed through a dryer before
106 entering the reactor. The difference between efflorescence RH (~35 %) and deliquescence RH (~80 %) for
107 AS aerosol is fairly large (Seinfeld and Pandis, 2016). Taking advantage of this property, it is possible to
108 vary the amount of water in aerosol under the same RH in the reactor. Initial seed number and volume
109 concentrations upon atomization for 20 min were approximately $2 \times 10^4 \text{ cm}^{-3}$ and $2 \times 10^{10} \text{ nm}^3 \text{ cm}^{-3}$,
110 respectively. A known amount of precursor VOC in the liquid form was transferred into a glass bulb, which
111 was then evaporated and carried into the reactor by flowing zero air at 5 L min^{-1} through the bulb. Although

112 the measurement of the precursor VOC concentration was not available for all experiments, the target and
113 measured concentrations in the experiments when the measurements were available were consistent.

114 For “daytime” experiments, an oxidant precursor (i.e., H₂O₂) was introduced to the reactor in the
115 same manner as VOC except that the glass bulb was gently heated by a heat gun to help evaporate faster.
116 During the injection of H₂O₂, a desired amount of NO was introduced into the reactor from a cylinder
117 containing 500 ppm NO (Matheson). For Exp. 3-5, 5 ppm NO at 5 L min⁻¹ was continuously injected to the
118 reactor until the SOA growth ceased. For Exp. 1, 2, 6, and 7, 15 ppm NO at 5 L min⁻¹ was injected for 5-
119 20 min several times until the SOA growth ceased. The NO concentration was usually on the order of tens
120 of ppb and always remained above a few ppb, making the bimolecular reaction with NO a favorable RO₂
121 reaction pathway. The experiment was initiated by turning on the irradiation of UV light approximately 20
122 min after the end of the last injection to ensure that particles and vapors were mixed well inside the reactor.

123 The procedure for “nighttime” experiments was the same until the end of the precursor VOC
124 injection. An oxidant precursor (i.e., N₂O₅) was pre-made in a flow tube by simultaneously injecting 500
125 ppm NO₂ (Matheson) at 0.4 L min⁻¹ and ~250 ppm O₃ (Jelight 610) at 0.5 L min⁻¹. A simple kinetic box
126 model was used to adjust the concentration of O₃ and flow rates of both NO₂ and O₃ to maximize the
127 production of N₂O₅ and minimize the concentration of O₃, such that the VOC was dominantly oxidized by
128 ~~NO₃ nitrate radicals~~. Once N₂O₅ entered the reactor, it thermally decomposed to generate NO₂ and
129 ~~NO₃ nitrate radicals~~. VOC was usually depleted within the first 15 min of the experiment. For Exp. 8 and
130 14, the injection order of precursor VOC and oxidant precursor was switched such that the injection of
131 VOC marked the beginning of the experiment. For Exp. 12 and 13, 25 ppm formaldehyde was added to the
132 reactor to enhance the branching ratio of RO₂⁻+HO₂⁻ (Schwantes et al., 2015; Boyd et al., 2015) by injecting
133 an appropriate amount of formalin solution (37 % HCHO, Sigma-Aldrich) in the same manner as the
134 injection of H₂O₂. We do not discuss the details of the effect of the injection order nor the effects of the
135 RO₂⁻ fate here as they had negligible impact on the results concerning hydrolysis.

136

137 2.2. Instrumentation and data analysis

138 A High Resolution Time-of-Flight Aerosol Mass Spectrometer (HR-ToF-AMS; Aerodyne
139 Research Inc.) measured the concentrations of non-refractory organics (Org), sulfate (SO₄), nitrate (NO₃),
140 ammonium (NH₄), and chloride (Chl) (DeCarlo et al., 2006). The data were analyzed using PIKA v1.16I
141 and the unity collection efficiency was applied to all datasets. For the majority of nitrate-containing aerosol
142 regardless of the class (i.e., inorganic or organic), the nitrate moiety (i.e., -NO₂, -ONO₂, and -O₂NO₂) was
143 known to be thermally fragmented into NO⁺ and NO₂⁺ and was measured as NO₃ (Farmer et al., 2010). As
144 many past studies have demonstrated the feasibility to separate the contribution of inorganic (NO_{3,Inorg}) and
145 organic nitrate (NO_{3,Org}) to the measured NO₃ based on the ratio of NO⁺ and NO₂⁺ (Fry et al., 2009; Farmer
146 et al., 2010; Xu et al., 2015b; Kiendler-Scharr et al., 2016; Fry et al., 2018), we used Eq. (1) presented in
147 Farmer et al (2010) to obtain NO_{3,Org}. R_{AN} (i.e., NO⁺/NO₂⁺ from ammonium nitrate) was obtained from the
148 routine ionization efficiency calibration of HR-ToF-AMS using 300 nm ammonium nitrate particles. The
149 drawback of this method is that R_{ON} (i.e., NO⁺/NO₂⁺ from organic nitrate aerosol) could vary depending on
150 the chemical composition (Xu et al., 2015b). In addition, a non-negligible contribution of ammonium nitrate
151 could be expected in experiments with deliquesced seed aerosol owing to high solubility of HNO₃. Thus,
152 we obtained the R_{ON} measured in low RH experiments for each system of VOC and oxidation condition. In
153 order to account for changes in R_{AN} over time, R_{ON} was scaled accordingly assuming that the ratio of R_{ON}
154 to R_{AN} in the same system was constant (Fry et al., 2013). R_{AN}, R_{ON}, and R_{ON}/R_{AN} values obtained in this
155 study were consistent with previously reported values (Fry et al., 2009; Bruns et al., 2010; Boyd et al.,
156 2015; Nah et al., 2016) and are summarized in Table S1.

157 A Filter Inlet for Gases and AEROsols (FIGAERO) (Lopez-Hilfiker et al., 2014) coupled to a High
158 Resolution Time-of-Flight Iodide Chemical Ionization Mass Spectrometer (HR-ToF-I-CIMS; Aerodyne
159 Research Inc.) detected a suite of gaseous and particulate oxidized organic species as well as selected
160 inorganic species (Bertram et al., 2011; Lee et al., 2014). The operation of FIGAERO-HR-ToF-I-CIMS
161 was detailed in the previous studies (Nah et al., 2016; Sanchez et al., 2016). Reagent ions were generated
162 by flowing a mixture of CH₃I and dry ultra high purity (UHP) N₂ (Airgas) through a polonium-210 source
163 (NRD; Model P-2021). The instrument measured gaseous compounds by sampling air from the reactor at

164 ~1.7 L min⁻¹ while collecting particles onto a Teflon filter. Upon completion of the collection period,
165 collected particles were desorbed by temperature-programmed dry UHP N₂ flow and subsequently analyzed
166 by HR-ToF-I-CIMS. Sensitivity could decrease if the amount of reagent ions were significantly depleted
167 and/or if the secondary chemistry in the ion-molecule reaction (IMR) chamber occurred at a significant
168 degree (Lee et al., 2014). To avoid changes in sensitivity among experiments, gas-phase sampling flow was
169 diluted with zero air immediately before the inlet such that the evaporation of aerosol was minimal. The
170 amount of aerosol collected on the filter was also adjusted by varying the sampling rate from 1 to 6 L min⁻¹
171 depending on the aerosol mass concentrations inside the reactor. Overall, the fraction of reagent ions to
172 the total ions was kept above 80 %. In addition, iodide ion chemistry has been known to be affected by the
173 water vapor pressure inside the IMR owing to the difference in thermodynamics between I⁻ and IH₂O⁻ to
174 analyte compounds (Lee et al., 2014). In order to minimize changes in the water vapor pressure inside the
175 IMR, a small continuous flow of humidified UHP N₂ (30-50 ccm) through a bubbler at a reduced pressure
176 was continuously added to the IMR directly. Therefore, while the instrument was not calibrated to report
177 the concentration of detected species, it was possible to quantitatively compare measured signal of each ion
178 among experiments because the instrument was operated under configurations that prevented undesired
179 changes in sensitivity. The data were analyzed using Tofware v2.5.11 and all the masses presented in this
180 study were I⁻ adducts.

181 A Scanning Mobility Particle Sizer (SMPS) that consisted of a Differential Mobility Analyzer (TSI
182 3040) and a ~~Ceondensation~~-Condensation Particle Counter (TSI 3775) was operated under the low flow
183 mode with the sheath flow of 2 L min⁻¹ to detect particles up to 1 μm in electrical mobility size. A Cavity
184 Attenuated Phase Shift spectroscopy (CAPS; Aerodyne Research Inc.) (Kebabian et al., 2005), an ultra-
185 sensitivity NO_x analyzer (Teledyne M200EU), and an UV absorption O₃ analyzer (Teledyne T400)
186 measured NO₂, NO_x, and O₃, respectively. In selected experiments, a Gas Chromatograph coupled to Flame
187 Ionization Detector (GC-FID; Agilent) was used to make sure that a desired amount of a precursor VOC
188 was injected. Except CAPS, NO_x analyzer, and O₃ analyzer, all instruments had their own dedicated
189 sampling line.

190

191 3. Results and discussion

192 3.1. Chemical composition of secondary organic aerosol

193 Shown in Fig. 1 are the mass spectra of particle-phase species obtained from FIGAERO-HR-ToF-
194 I-CIMS at peak SOA growth. Many of the major species detected in this study are previously reported using
195 the same or different technique (Eddingsaas et al., 2012; Claflin and Ziemann, 2018; Boyd et al., 2015; Nah
196 et al., 2016; Lee et al., 2016; Romonosky et al., 2017). Concerning the chemical composition of SOA from
197 each system, a more distinct difference is observed between different oxidation conditions (i.e.,
198 ~~OH· hydroxyl~~-vs. ~~NO₃· nitrate radical~~-oxidation) than between different precursor VOC (i.e., α -pinene vs.
199 β -pinene). This is expected as α -pinene and β -pinene have the same chemical formula and very similar
200 structures while the oxidation condition is distinctively different (Kroll and Seinfeld, 2008; Ziemann and
201 Atkinson, 2012).

202 For “daytime” experiments where ~~OH· hydroxyl radicals~~ are the dominant oxidants, the
203 contribution of ON (i.e., $C_xH_yN_{1,2}O_z$) and non-nitrated organics (i.e., $C_xH_yO_z$) are comparable and their
204 contributions are well spread out over a wide range of masses. A large contribution of $C_xH_yO_z$ is expected
205 because the formation of ON is a minor pathway in $RO_2 + NO$ (Perring et al., 2013). In Eddingsaas et al.
206 (2012), the major compounds reported in the α -pinene+OH· system include $C_8H_{12}O_{4,6}$ and $C_{10}H_{16}O_{4,6}$,
207 which are also detected in our study. A suite of C_{10} ON from the chamber experiment of the α -pinene+OH·
208 system are reported in Lee et al. (2016) with the chemical formula of $C_{10}H_{15,17,19}NO_{4-9}$. All of these masses
209 are detected in this study, though we observe a considerable contribution of $C_{<10}$ ON (i.e., $C_7H_9,11NO_6$ and
210 $C_9H_{13,15,17}NO_6$) as well as a small fraction of C_{10} dinitrate (i.e., $C_{10}H_{14,16}N_2O_{9,10}$) that has been rarely reported
211 (Fig. 1a). This significant contribution from species containing $C_{<10}$ indicates the large contribution of
212 fragmentation process that is a preferred pathway in high NO conditions (Kroll and Seinfeld, 2008;
213 Ziemann and Atkinson, 2012; Perring et al., 2013). It is possible that these species with $C_{<10}$ are thermally
214 decomposed products during the thermal desorption process (Stark et al., 2017). However, it is unlikely

215 that thermal decomposition plays a significant role for the SOA generated via OH· hydroxyl radical
216 oxidation of monoterpenes because the desorption temperature for these compounds (i.e., peaking at ~50-
217 70 °C) ~~are-is~~ much lower than the temperature at which decarboxylation or dehydration reactions (>120
218 °C) are expected to occur (Stark et al., 2017). SOA chemical composition of the β -pinene+OH· system is
219 similar to that of the α -pinene+OH· system but with a larger contribution of $C_xH_yO_z$.

220 Another interesting observation in the α -/ β -pinene+OH· systems is that a selected class of
221 compounds (i.e., $C_{10}H_{13,15,17}NO_{5-8}$) with the same H/C exhibit the same ~~the~~ time evolutions regardless of
222 the number of oxygen (Fig. S1). This observation is consistent with the autoxidation mechanism, in which
223 highly oxidized molecules are formed in a short time scale (Ehn et al., 2014; Crounse et al., 2013; Jokinen
224 et al., 2015). Based on FIGAERO-HR-ToF-I-CIMS data (Fig. S1), $C_{10}H_{17}NO_{>6}$ peak at 75 min. This is
225 comparable to the lifetime of α -pinene at 53 min in the same experiment, which is related to characteristic
226 time of OH oxidation. This suggests that the aforementioned ON are likely formed via one OH oxidation
227 reaction, which is consistent with the autoxidation scheme to generate $C_{10}H_{17}NO_{>6}$ proposed in prior studies
228 (Berndt et al., 2016; Xu et al., 2019; Pye et al., 2019). The detection of compounds with $O_{>7}$ also supports
229 the presence of autoxidation process; otherwise it is extremely difficult to produce highly oxidized
230 molecules in a timescale of a few hydroxyl radical oxidation reactions. It is important to note that the
231 concentration of NO in our experiments has been mostly kept on the order of tens of ppb over the course
232 of the experiments by a continuous injection of dilute NO and, therefore, this result suggests that
233 autoxidation is not a negligible pathway of RO_2 · fate even at a moderately high NO level in laboratory
234 experiments and in polluted ambient environments. Indeed, recent studies (Berndt et al., 2016; Xu et al.,
235 2019; Pye et al., 2019) suggest that the autoxidation rate constant for the α -pinene+OH· system could be
236 up to a few per second, which is comparable to the NO level of ~10 ppb, assuming a typical RO_2 ·+NO
237 reaction rate constant of $1 \times 10^{-11} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ (Orlando and Tyndall, 2012). The autoxidation rate
238 constant as well as the role of NO in autoxidation based on this observation will be discussed in details in
239 a forthcoming publication.

240 In contrast, the signals of $C_xH_yN_{1-2}O_z$ are dominant in the ~~$NO_3\cdot$ nitrate radical~~-oxidation condition,
241 indicating that the production of ON is greatly favored over non-nitrated organics. This observation is
242 consistent with a direct addition of a nitrate functional group to a double bond (Wayne et al., 1991; Ng et
243 al., 2017), whereas the formation of ON in ~~the hydroxyl radical~~ $OH\cdot$ oxidation condition is a minor channel
244 of $RO_2\cdot+NO$ reaction (Perring et al., 2013). Although this is generally true for many monoterpenes, the
245 organic nitrate yield in α -pinene+ $NO_3\cdot$ has been known to be low (Fry et al., 2014) due to loss of a nitrate
246 functional group followed by alkoxy radical bond scission (Kurten et al., 2017). Fry et al. (2014) observe
247 no SOA formation in the same system, though another study has reported the formation of non-negligible
248 SOA mass even at a relatively low initial concentration of α -pinene (Nah et al., 2016). Since the particle-
249 phase compounds represent <10% of overall α -pinene+ $NO_3\cdot$ products by mass, it is not necessarily
250 inconsistent to observe a higher abundance of ON than non-nitrated organics in the particle phase if low-
251 volatility compounds mainly consist of ON.

252 Moreover, the contribution from species containing $C_{<10}$ is minimal in the $NO_3\cdot$ oxidation
253 condition. This Once $NO_3\cdot$ attacks a double bond in the initial oxidation reaction, the majority of the
254 reaction products no longer contain any double bond. is because u Unlike hydroxyl radical $OH\cdot$ -oxidation, a
255 hydrogen ~~subtraction~~-abstraction reaction by ~~a nitrate radical~~ $NO_3\cdot$ is ~~less efficient~~lower by orders of
256 magnitude (Atkinson and Arey, 2003). and, thus Therefore, multi-generation oxidation is unlikely to occur
257 within the timescale of experiments (Wayne et al., 1991; Ng et al., 2017). This means that once the precursor
258 VOC undergoes functionalization upon the initial ~~$NO_3\cdot$ nitrate radical~~-oxidation, it is not likely to
259 experience fragmentation during the experiment. Unlike “daytime” SOA, the distribution of masses is
260 dominated by a few signature ions, such as $C_{10}H_{15}NO_{5,6}$ and $C_{20}H_{32}N_2O_{8-10}$ (Fig. 1c and 1d). In Boyd et al.
261 (2015), Nah et al. (2016), and Lee et al. (2016), the major species reported in the α -/ β -pinene+ $NO_3\cdot$ systems
262 are monomeric nitrate aerosol (i.e., $C_{10}H_{13,15,17,19}NO_{4-10}$), while in this study a substantial contribution of
263 dimeric species (i.e., $C_{20}H_{32}N_2O_{8-11}$) is observed. The abundant presence of dimeric compounds has been
264 previously observed in some studies on ~~$NO_3\cdot$ nitrate radical~~-oxidation of α -/ β -pinene (Romonosky et al.,

265 2017; Clafin and Ziemann, 2018) and particle-phase reactions to produce such dimers have been proposed
266 by Clafin and Ziemann (2018). Many of the reported species in Clafin and Ziemann (2018) except the
267 trimeric species (mass scan range not extended to trimeric species in our study) are observed in our study.
268 One major difference between Clafin and Ziemann (2018) and this study is the substantial presence of
269 monomeric nitrate aerosol (i.e., 30-60 % by signal) in this study. This difference may be attributed to the
270 difference in the amount of available monomeric blocks to form dimer species. Assuming reversible
271 dimerization process, the concentration of dimer species shall be proportional to the square of monomer
272 concentration, such that the monomer to dimer ratio increases in a quadratic manner as the available
273 monomer concentration decreases. Since the amount of SOA formed in Clafin and Ziemann (2018) is
274 approximately two orders of magnitude higher than that in this study, the concentration of monomers in the
275 particle-phase is higher, favoring a more efficient formation of dimeric species. Together, these results
276 suggest that the contribution of dimeric nitrate aerosol could vary greatly depending on the concentrations
277 of monomeric blocks at the specific time and location.

278 Previous field studies have reported the mass spectra of ambient C_{10} pON obtained by FIGAERO-
279 HR-ToF-I-CIMS in rural Alabama site during the Southern Oxidant and Aerosol Study (SOAS) (Lee et al.,
280 2016) and in rural forest in Germany (Zhang et al., 2018). A comparison of the ambient mass spectra with
281 those obtained in this study reveals that average ambient pON resembles “daytime” pON more than
282 “nighttime” pON (Fig. S2). pON from “daytime” experiments has a distribution of masses centered around
283 $C_{10}H_{13,15,17}NO_7$, which is consistent with the ambient measurement data. On the other hand, ~~NO₃· nitrate~~
284 ~~radical~~ oxidation does not seem to oxidize organic species enough that the distribution of masses is skewed
285 towards a less-oxidized region (i.e., $C_{10}H_{13,15,17}NO_{5-6}$). However, it is difficult to draw a quantitative
286 conclusion simply based on this comparison because O₃, another important oxidant at night, is not studied.
287 Moreover, an average lifetime of aerosol could extend up to a week and thus ambient aerosol is continuously
288 exposed to further oxidation while the experiments here are more applicable to freshly formed aerosol. In
289 addition, the use of C_{10} pON alone may not be a good representative of monoterpene-derived pON as 42-74
290 % of pON in this study contains fewer or more than 10 carbon (Table S2). Nonetheless, the chemical

291 composition of aerosol generated in this study is comparable to those in the atmosphere and, thus, the results
292 shall be directly applicable to relevant ambient conditions.

293

294 3.2. Hydrolysis of particulate organic nitrate

295 3.2.1. Proxy used to evaluate hydrolysis process

296 Various proxies using HR-ToF-AMS data have been used to infer ON hydrolysis in previous
297 studies. In Bean and Hildebrandt Ruiz (2016), NO_3 measured by an Aerosol Chemical Speciation Monitor
298 (ACSM, practically similar to how AMS operates and measures aerosol species) (Ng et al., 2011) is
299 normalized to SO_4 as a means to account for particle wall-loss and is fitted by an exponential function to
300 estimate the ON hydrolysis rate. On the other hand, Boyd et al. (2015) normalize NO_3 measured by HR-
301 ToF-AMS to Org and attribute the relative decay of NO_3/Org between humid (RH ~50 %) to dry (RH <5
302 %) conditions to hydrolysis. Other approaches include the SMPS-derived particle wall-loss correction of
303 NO_3 measured by HR-ToF-AMS followed by fitting its decay trend (Liu et al., 2012b) and the determination
304 of fraction of total (i.e., gas and particle) ON to the precursor VOC consumed as a function of RH using a
305 Fourier-Transform InfraRed spectroscopy (FTIR). Below, we systematically examine the use of different
306 proxies using HR-ToF-AMS data to infer hydrolysis and discuss how the corresponding results are
307 interpreted.

308 Figure 2 shows the time series of Org, NO_3 , SO_4 , and NH_4 measured by HR-TOF-AMS for α -
309 pinene+ NO_3 system. There is a substantial difference in NO_3 for the same VOC system but under different
310 reactor RH and phase state of seed particles (Exp. 10 and 11), while Org and SO_4 concentrations are similar.
311 The spike in NO_3 in high RH wet seed experiment (Fig. 2b) is attributed to the uptake of N_2O_5 and/or
312 dissolution of HNO_3 into aqueous aerosol followed by neutralization with ammonia to produce ammonium
313 nitrate. This is consistent with the sharp increase in molar ratio of NH_4/SO_4 to higher than 2, which is the
314 theoretical value for AS particles. It is also possible that inorganic nitrate is generated via hydrolysis of
315 gaseous ON that is too volatile to condense but is soluble enough to dissolve in aqueous aerosol and, thus,
316 only appears in high RH experiments. Since we do not have a way to quantitatively differentiate the

317 contribution of the aforementioned sources, the focus of this study is on hydrolysis of p ON that ~~is-partitions~~
318 ~~to the aerosol due to condensation rather than dissolution, condensed-driven-by-volatility~~. However, further
319 study is required to investigate the hydrolysis of volatile yet soluble gaseous ON, and the approach must be
320 different from the comparison between low and high RH experiments to obtain meaningful results.

321 To evaluate the extent of p ON hydrolysis, the contributions of inorganic nitrate ($\text{NO}_{3,\text{Inorg}}$) and
322 organic nitrate ($\text{NO}_{3,\text{Org}}$) to the measured NO_3 need to be calculated. Firstly, $\text{NO}_{3,\text{Org}}$ is estimated by
323 subtracting NO_3 associated with excess NH_4 . Secondly, $\text{NO}_{3,\text{Org}}$ is derived from $\text{NO}^+/\text{NO}_2^+$ approach (Sect.
324 2.2.). Figure S3 shows the comparison of $\text{NO}_{3,\text{Org}}$ estimated by these two independent methods for the α -
325 pinene+ NO_3 system. It is clear that $\text{NO}_{3,\text{Org}}$ from both methods are consistent and that there is a
326 considerable contribution of $\text{NO}_{3,\text{Inorg}}$ to NO_3 in the experiment. We note the contribution of $\text{NO}_{3,\text{Inorg}}$ to
327 NO_3 can vary depending on experimental conditions with a range from 28 to 90 % for all experiments in
328 this study. Nevertheless, these results demonstrate that for laboratory experiments with high RH and wet
329 seeds, when using HR-ToF-AMS data to infer hydrolysis, it is important to separate the measured NO_3 into
330 $\text{NO}_{3,\text{Inorg}}$ and $\text{NO}_{3,\text{Org}}$.

331 Once NO_3 is separated into $\text{NO}_{3,\text{Inorg}}$ and $\text{NO}_{3,\text{Org}}$, we evaluate whether the normalization of $\text{NO}_{3,\text{Org}}$
332 to SO_4 and Org provides a consistent decay trend. Photooxidation of α -pinene (Exp. 3-5) is used as a case
333 system. As hydrolysis is a reaction ~~where-in-which~~ liquid water is ~~directly-involved-in-a-reactant~~, it is
334 expected that the rate of hydrolysis will change proportionally as a function of aerosol water content. Based
335 on the hygroscopicity parameter for AS ($\kappa = 0.53$) (Petters and Kreidenweis, 2007) and for ambient LO-
336 OOA ($\kappa = 0.08$) (Cerully et al., 2015) that has a substantial contribution from p ON (Xu et al., 2015a),
337 estimated aerosol water contents at peak SOA growth in Exp. 3-5 are approximately 0, 1, and 26 $\mu\text{g m}^{-3}$,
338 respectively. Figure S4 illustrates that AS seed particles are indeed effloresced in Exp. 4 (high RH, dry AS)
339 but not in Exp. 5 (high RH, wet AS). These mass concentrations of aerosol water translate to 0, 6, and 36
340 mol L^{-1} , respectively, under the assumption that SOA is miscible with liquid water. It is speculated that
341 SOA formed in Exp. 3-5 are miscible with water because (1) the measured O/C ratio in HR-ToF-AMS

342 (Canagaratna et al., 2015) is close to 0.7, which is near the lower end but above the liquid-liquid phase
343 separation condition (Song et al., 2012) and (2) there is evidence of aqueous-phase reactions which highly
344 depend on the availability of aerosol water, as discussed in Sect. 3.3. Thus, the decay rate of $\text{NO}_{3,\text{Org}}$
345 normalized to SO_4 and/or Org between Exp. 4 and Exp. 5 shall differ by a factor of 6 based on the molar
346 concentrations of aerosol water.

347 Figure 3a shows the mass ratio of $\text{NO}_{3,\text{Org}}/\text{SO}_4$ and the decay rate as a characteristic time in Exp. 3-
348 5. The characteristic times of Exp. 4 and 5 (4.4 vs. 4.0 h) do not differ regardless of the molar concentrations
349 of aerosol water, suggesting that the decreasing trend in $\text{NO}_{3,\text{Org}}/\text{SO}_4$ may not be due to changes in aerosol
350 water content and pON hydrolysis, but arise from the difference in the reactor humidity alone. A comparison
351 of $\text{NO}_{3,\text{Org}}/\text{SO}_4$ and $\text{NO}_{3,\text{Org}}/\text{Org}$ also reveals that these two proxies capture a different range of decay
352 mechanisms. Figure 3b shows the relative decay trend of $\text{NO}_{3,\text{Org}}/\text{SO}_4$ and $\text{NO}_{3,\text{Org}}/\text{Org}$ between Exp. 4 (high
353 RH) and Exp. 3 (low RH). If hydrolysis is a dominant decay mechanism of pON , the trend of $\text{NO}_{3,\text{Org}}/\text{Org}$
354 would be identical to that of $\text{NO}_{3,\text{Org}}/\text{SO}_4$. This is because the organic moiety of hydrolysis product is
355 generally considered to have a substituted alcohol group (Boschan et al., 1955) and to have a relatively
356 similar vapor pressure and shall remain in the particle phase (Pankow and Asher, 2008). Similar to this is
357 the formation of organic sulfate from hydrolysis of pON (Liggio and Li, 2006, 2008; Surratt et al., 2008),
358 which has sufficiently low volatility to remain in the particle phase as with alcohol substituted products.
359 However, the measured decay trend of the two proxies is greatly different. It is possible that some organic
360 moiety of hydrolysis product could be significantly more volatile and repartition back to the gas phase
361 (Rindelaub et al., 2016; Bean and Hildebrandt Ruiz, 2016) and, thus, both organics and HNO_3 formed from
362 hydrolysis evaporate. In this case, not only $\text{NO}_{3,\text{Org}}$ but also some fraction of Org would decrease because
363 Org measured by HR-ToF-AMS includes the contribution from the organic part of pON . This will lead to
364 the relatively smaller decrease in $\text{NO}_{3,\text{Org}}/\text{Org}$ compared to $\text{NO}_{3,\text{Org}}/\text{SO}_4$. We can reconstruct the decay rate
365 of $\text{NO}_{3,\text{Org}}/\text{Org}$ assuming 1) the decay rate of $\text{NO}_{3,\text{Org}}/\text{SO}_4$ is solely due to hydrolysis of pON and 2) the
366 maximum contribution of pON to OA is 35 % (see Fig. 4 and discussions below). The reconstructed decay
367 rate of $\text{NO}_{3,\text{Org}}/\text{Org}$ is shown in Fig. 3b. As observed in the figure, the decay rate of the reconstructed

368 $NO_{3,Org}/Org$ is much larger than the measured $NO_{3,Org}/Org$. This suggests that hydrolysis is not the only loss
369 process reflected in the decreasing trend of $NO_{3,Org}/SO_4$, while $NO_{3,Org}/Org$ is likely a better proxy that
370 isolates hydrolysis from other loss processes. The likely important loss process manifested in $NO_{3,Org}/SO_4$
371 is the loss of organic vapors to the walls of the ~~reactor~~ reactor (Matsunaga and Ziemann, 2010; Krechmer et
372 al., 2016; Huang et al., 2018; Loza et al., 2010; Mcvay et al., 2014; Zhang et al., 2015; Zhang et al., 2014;
373 La et al., 2016). For example, Huang et al. (2018) observe that the decay of isoprene hydroxyl nitrate
374 depends on the reactor humidity. While SO_4 is practically non-volatile in the experimental condition of this
375 study, both ON and non-nitrated organics could have some fractions of semi-volatile species whose vapors
376 are subject to wall loss. Assuming a uniform loss rate of ON and non-nitrated organic vapors to the reactor
377 walls, the effect of vapor wall loss could be effectively cancelled out in $NO_{3,Org}/Org$, but not in $NO_{3,Org}/SO_4$.
378 This assumption is likely reasonable because the vapor wall loss rate is a function of saturation mass
379 concentration (Zhang et al., 2015) and the average saturation mass concentrations of bulk ON and non-
380 nitrated organic aerosol are similar based on the thermal desorption profiles in FIGAERO-HR-ToF-I-CIMS
381 (Fig. S5). Therefore, $NO_{3,Org}/Org$ is a better proxy to infer hydrolysis of pON than others.

382

383 3.2.2. Hydrolysis lifetime of particulate organic nitrate

384 In order for the data to be easily comparable with those reported in models or using other
385 techniques, the use of general terms instead of the AMS specific terms (i.e., $NO_{3,Org}$ and Org) can be
386 convenient. We define pON as the total mass concentration of particulate organic nitrate (includes organics
387 part and nitrate part of the ON compounds) and OA as the total mass concentration of organic aerosol
388 (includes nitrate and non-nitrated organics). The inclusion of nitrate mass concentration in OA is important
389 as the contribution of nitrate functional groups to the total organic aerosol mass concentration is large. The
390 conversion method from $NO_{3,Org}/Org$ into pON/OA is illustrated in Eq. (1).

$$\frac{pON}{OA} = \left(\frac{NO_{3,Org}}{Org + NO_{3,Org}} \right) \times \left(\frac{MW_{pON}}{MW_{NO_3}} \right) = \left(\frac{\frac{NO_{3,Org}}{Org}}{1 + \frac{NO_{3,Org}}{Org}} \right) \times \left(\frac{MW_{pON}}{MW_{NO_3}} \right) \quad (1)$$

391 MW_{pON} represents the average molecular weight of pON per nitrate functional group estimated from
392 FIGAERO-HR-ToF-I-CIMS data assuming a uniform sensitivity among detected species, and MW_{NO₃}
393 indicates the molecular weight of nitrate (i.e., 62 g mol⁻¹). Since MW_{pON} does not significantly vary during
394 the course of experiments (i.e., relative standard deviation of <1.2 %), the average value is applied to each
395 experiment. The variability of MW_{pON} among different systems is also found to be small, ranging from 229
396 to 238 g mol⁻¹ and, thus, an average MW_{pON} is used for experiments where FIGAERO-HR-ToF-I-CIMS
397 data are not available. Figure 4 shows the time-series data of pON/OA for all the systems investigated in
398 this study.

399 For “nighttime” experiments, the relative ratio of C_xH_yN₁₋₂O_z and C_xH_yO_z obtained from
400 FIGAERO-HR-ToF-I-CIMS data in Fig. 1 does not appear to match well with pON/OA from HR-ToF-
401 AMS data. For example, the signals are dominated by C_xH_yN₁₋₂O_z in the β-pinene+NO₃ system, as shown
402 in Fig. 1d, while pON/OA is at most 0.5, as shown in Fig. 4d. The discrepancy would be attributed to the
403 overestimation of Org (in particular, C_xH_y family) in HR-ToF-AMS and/or underestimation of C_xH_yO_z in
404 FIGAERO-HR-ToF-I-CIMS. Relative ionization efficiency (RIE) of less-oxidized organic species in HR-
405 ToF-AMS is experimentally measured to be higher at least by a factor of 2 (Xu et al., 2018b). As previously
406 reported (Boyd et al., 2015), the HR-ToF-AMS mass spectrum of SOA formed from β-pinene+NO₃
407 contains significant amounts of C_xH_y fragments, indicating the less-oxidized nature of SOA from β-
408 pinene+NO₃. For example, if the true RIE of Org by β-pinene+NO₃ SOA were to be a factor of 2 higher
409 than the default RIE of Org (i.e., 1.4), the reported concentration of Org would have been overestimated by
410 a factor of 2, such that actual pON/OA would have been higher than reported in Fig. 4d. On the other hand,
411 an iodide reagent ion is not quite selective to detect less oxidized species, which could overestimate the
412 contribution of pON to OA (Aljawhary et al., 2013). Nonetheless, this discrepancy between HR-ToF-AMS
413 and FIGAERO-HR-ToF-I-CIMS data, however, should not affect the results regarding the hydrolysis
414 lifetime and hydrolyzable fraction of pON presented later.

415 As illustrated in Fig. 4, the time series of pON/OA stabilizes fairly quickly in most of the
416 experiments, regardless of RH and/or the phase state of seed aerosol, supporting the appropriateness of the

417 reconstructive approach shown in Fig. 3b. This suggests that the timescale of p ON hydrolysis in the systems
418 studied is significantly shorter or longer than the timescale of our experiments. It is also evident from Fig.
419 4 that p ON/OA in high RH experiments are always lower than that in low RH experiments. These two
420 observations imply that the rate of p ON hydrolysis may be fast enough that the decay trend of p ON compared
421 to OA is not visibly manifested, though a clear difference of p ON/OA between low and high RH
422 experiments is exhibited as a result of fast hydrolysis. Since no sudden, drastic change in p ON/OA is
423 observed except for a few initial data points, we conclude that the hydrolysis lifetime of hydrolyzable p ON
424 for α -/ β -pinene derived ON shall be no more than 30 min (i.e., 3 data points in Fig. 4). Particle acidity is
425 found to enhance ~~rate of hydrolysis in prior study~~hydrolysis rate of α -pinene hydroxy nitrate (Rindelaub et
426 al., 2016), though no clear difference is observed here between experiments with AS and SA+MS seed
427 particles (i.e., Exp. 5 and 2). It is worth noting that aqueous AS particles are not neutral but slightly acidic
428 due to partitioning of ammonium to the gas phase once the particles enter the chamber (Gao et al., 2004).
429 In Gao et al. (2004), the reported pH of aqueous AS particles is 4.6 and we expect a similar pH in our study.
430 In Rindelaub et al. (2016), the reported hydrolysis lifetime is short at 1.3 h at pH of 4.0. On the other hand,
431 previous studies have shown that isoprene-derived hydroxy nitrates do not require low pH to undergo fast
432 hydrolysis (Darer et al., 2011; Hu et al., 2011). Thus, p ON formed from OH· hydroxyl and NO₃· nitrate
433 ~~radical~~ oxidation of α -pinene and β -pinene may not require a low pH to undergo hydrolysis at a rate
434 comparable to the timescale of chamber experiments.

435 Comparing with the results from past chamber studies reporting a p ON hydrolysis lifetime of 3 to
436 6 h (Liu et al., 2012b; Bean and Hildebrandt Ruiz, 2016; Boyd et al., 2015), our estimated hydrolysis
437 lifetime is substantially shorter, but is consistent with the range (i.e., 1 min to 8.8 h) reported in studies
438 using the bulk solution method (Darer et al., 2011; Hu et al., 2011; Jacobs et al., 2014; Rindelaub et al.,
439 2016) (Fig. 5). In Liu et al. (2012) and Bean and Hildebrandt Ruiz (2016), the hydrolysis lifetime is derived
440 from the decay rate of NO₃ corrected for the particle wall loss or normalized to SO₄. As demonstrated in
441 Sect. 3.2.1., this NO₃/SO₄ decay rate is likely affected by other loss processes, such as vapor wall loss, and,
442 thus, is not a good proxy to estimate the hydrolysis lifetime. The apparent discrepancy does not stem from

443 the contradiction in the obtained data itself but rather in the data interpretation. Indeed, the lifetime
444 estimated based on the decay of $\text{NO}_{3,\text{Org}}/\text{SO}_4$ in our study is 4 h (Fig. 3a) that is consistent with 6 and 3 h
445 reported in Liu et al. (2012) and Bean and Hildebrandt Ruiz (2016). On the other hand, in Boyd et al. (2015)
446 the hydrolysis lifetime is estimated based on the decay rate of NO_3/Org . The discrepancy in the reported
447 hydrolysis lifetime here could be attributed to the fact that NO_3 is not separated into $\text{NO}_{3,\text{Org}}$ and $\text{NO}_{3,\text{Inorg}}$.
448 Figure [S5-S6](#) shows our data analyzed in the same manner as Boyd et al. (2015). The lifetime calculated
449 based on the decay is 2.2 h, which is close to the reported [3-4.5](#) h (Boyd et al., 2015). The reduction of
450 NO_3/Org in Fig. [S5-S6](#) (~30 %) is greater than in Boyd et al. (2015) (10 %), which could be because (1)
451 the amount of N_2O_5 , a source of inorganic nitrate, used in our study is slightly larger and (2) RH in our
452 study is higher by 10-20 %, which may have allowed greater uptake of N_2O_5 to produce inorganic nitrate
453 [due to changes in aerosol viscosity](#) (Grzanic et al., 2015).

454 In previous bulk solution studies where the concentration of interested organic nitrate (in particular
455 hydroxy⁴ nitrate) in aqueous solution rather than in aerosol water is monitored over time, it is unlikely that
456 the data interpretation is affected by other loss processes present in chamber experiments, such as vapor
457 wall-loss. It is also common to monitor the organic moiety of hydrolysis product (Darer et al., 2011), while
458 it is extremely difficult in chamber experiments where hundreds of organic species are present at the same
459 time, leading to the difficulty in accurately measuring the hydrolysis lifetime in chamber experiments.
460 Based on the comprehensive analysis we demonstrate above on evaluating pON hydrolysis in chamber
461 experiments, we recommend the use of the hydrolysis lifetime reported in this study, which is no more than
462 30 min, for pON formed from α -pinene and β -pinene.

463

464 3.2.3. Hydrolyzable fraction of particulate organic nitrate

465 The fraction of hydrolyzable pON (F_H) can be directly estimated from the difference in pON/OA
466 between low and high RH experiments shown in Fig. 4. Although we show that hydrolysis is substantially
467 faster than the timescale of chamber experiments in our study, there still appears a clear difference in
468 pON/OA between high RH experiments but with a different phase state of seed aerosol (i.e., Exp. 4 and 5).

469 The difference mass spectra among Exp. 3-5 obtained from FIGAERO-HR-ToF-I-CIMS reveal that the
470 difference in p_{ON}/OA between Exp. 4 and 5 does not arise from the reduction in p_{ON} but from the increase
471 in non-nitrated organics (Fig. 6). The reason for this OA increase with the abundant presence of aerosol
472 water is speculated to be uptake and other aqueous-phase reactions than hydrolysis and is briefly discussed
473 in Sect. 3.3. Thus, the absolute difference in p_{ON}/OA between low RH, dry seed and high RH, dry seed
474 experiments best indicates F_{H} . Depending on the fate of the organic moiety of the hydrolysis product (i.e.,
475 stay in the particle phase or repartition back to the gas ~~phase~~), F_{H} varies. Since we are unable to
476 determine the fate of hydrolysis product, an upper and lower limit of F_{H} are reported as a range of F_{H} . For
477 the α -pinene+OH \cdot system, 23-32 % of p_{ON} formed undergoes hydrolysis within the timescale of the
478 experiments. For the other systems with no experiment under high RH with dry seed aerosol, the same level
479 of additional contribution from non-nitrated organics encountered in α -pinene+OH \cdot system is assumed and
480 F_{H} is scaled accordingly. For the β -pinene+OH \cdot , α -pinene+NO $_3\cdot$, and β -pinene+NO $_3\cdot$ systems, 27-34, 9-
481 17, and 9-15 % of p_{ON} are found hydrolyzable within the timescale of the experiments. Table 2 summarizes
482 the hydrolysis lifetime and F_{H} in the systems explored in this study.

483 F_{H} has been only reported in ~~several a few~~ studies (Liu et al., 2012b; Boyd et al., 2015; Zare et al.,
484 2018). The determination of F_{H} is essential because the assumption that all p_{ON} hydrolyzes biases the
485 relative importance of hydrolysis among the loss mechanisms of p_{ON} and NO $_x$. Boyd et al. (2015) report
486 that F_{H} of p_{ON} formed via β -pinene+NO $_3\cdot$ is ~ 10 %, which is in a good agreement with our range of 9-15
487 %. From a perspective of predicted molecular structures of p_{ON} , <5 % of p_{ON} from β -pinene+NO $_3\cdot$ are
488 tertiary (Claflin and Ziemann, 2018) that is expected to undergo hydrolysis in minutes (Darer et al., 2011;
489 Hu et al., 2011). In our study, a considerable contribution of monomeric (C $_{10}$) p_{ON} is observed (Fig. 1d),
490 while dimeric (C $_{20}$) p_{ON} is dominant in Claflin and Ziemann (2018). This may indicate that monomeric
491 p_{ON} is more susceptible to hydrolysis such that F_{H} in this study is slightly higher than expected based on
492 the proposed molecular structures of p_{ON} in Claflin and Ziemann (2018).

493 F_H for α -/ β -pinene+OH \cdot systems are higher than those from α -/ β -pinene+NO $_3\cdot$ systems. This trend
494 is consistent with the understanding that pON via OH \cdot hydroxyl radical oxidation have a larger fraction of
495 tertiary nitrate groups, which are significantly more susceptible to hydrolysis (Darer et al., 2011; Hu et al.,
496 2011) than those formed via NO $_3\cdot$ nitrate radical oxidation (Ng et al., 2017). Although the relative trend of
497 F_H between α -/ β -pinene+OH \cdot and α -/ β -pinene+NO $_3\cdot$ systems is consistent, the magnitude of F_H in the α -
498 / β -pinene+OH \cdot system appears to be smaller than expected based on the tertiary nitrate fraction estimated
499 via the explicit gas-phase chemical mechanism (Browne et al., 2013; Zare et al., 2018). In previous studies,
500 F_H are estimated to be 62 and 92 % for α -/ β -pinene+OH \cdot systems, respectively. However, our results
501 indicate F_H is 23-32 and 27-34 % for the same systems (Fig. 5). The chemical mechanism used in Browne
502 et al. (2013) and Zare et al. (2018) are based on the Master Chemical Mechanism (MCM) that is well known
503 in the degradation chemistry of VOC in the gas phase (Jenkin et al., 1997; Saunders et al., 2003). However,
504 the same mechanism performs poorly in regards to the chemical composition of SOA (Faxon et al., 2018)
505 as well as the prediction of SOA formation (Ruggeri et al., 2016; Boyd et al., 2017) when equipped with
506 gas-particle partitioning modules based on the absorptive gas-particle partitioning theory (Pankow, 1994).
507 It is, therefore, reasonable to argue that the chemical composition of pON could greatly differ from that of
508 total ON predicted by the MCM. Thus, F_H reported in this study provides the fundamental experimental
509 constraints on pON hydrolysis, that can be used in regional and global models for elucidating potential
510 impacts of ON on nitrogen budget and formation of ozone and aerosol.

511

512 3.3 Signature of other aqueous-phase reactions than hydrolysis

513 As briefly discussed in the above section, the presence of elevated level of aerosol water seems to
514 have enhanced the contribution of small (i.e., C $_9$) but more-oxidized organic species to SOA. As shown in
515 Fig. 5c, the enhancement is observed for a variety of non-nitrated organic aerosol including C $_4$ H $_{6,8}$ O $_3$,
516 C $_4$ H $_4$ O $_4$, C $_5$ H $_6$ O $_5$, C $_7$ H $_{8,10}$ O $_4$, and C $_8$ H $_{8,10}$ O $_{4,5}$, while pON overall neither increase nor decrease. C $_5$ H $_6$ O $_5$ has
517 been reported as a product of the aqueous-phase reaction of α -pinene derived compounds (Aljawhary et al.,
518 2016). Other compounds, such as C $_4$ H $_4$ O $_4$, C $_7$ H $_{8,10}$ O $_4$, and C $_8$ H $_{10}$ O $_5$, also appear to result from the aqueous

519 processing because compounds with similar chemical formulae but with slightly higher degrees of oxidation
520 (i.e., $C_4H_4O_5$, $C_7H_{10}O_5$, and $C_8H_{12}O_6$) are reported in Aljwhary et al. (2016). The reason for this less
521 oxidized nature of SOA in this study may be attributed to our experiments being performed in moderately
522 high NO conditions that promotes a higher contribution of a carbonyl group than a hydroperoxyl group,
523 which is preferred in low NO conditions. In Aljwhary et al. (2016), the starting compound is a product in
524 low NO conditions (pinonic acid, $C_{10}H_{16}O_3$). Thus, it is reasonable that products of the aqueous processing
525 in this study are slightly less oxidized than observed in Aljwhary et al. (2016).

526 The enhancement of non-nitrated organic aerosol in FIGAERO-HR-ToF-I-CIMS may be due to
527 aqueous processing of species in the particles in the presence of aerosol water. This can come from further
528 reactions of semi/low-volatile species in the particle phase, or reactive uptake of volatile (but highly water-
529 soluble) species into the aerosol followed by subsequent aqueous-phase reactions to form low-volatility
530 products (Ervens et al., 2011). A comparison of the average thermogram at peak SOA growth among Exp.
531 3-5 indicates a higher contribution of low-volatility compounds in Exp. 5 than in Exp. 3 and 4, as illustrated
532 by the bimodal peaks (Fig. S6S7). Given the same degree of gas-phase oxidation expected among Exp. 3-
533 5, these results show that that aqueous chemistry in wet aerosol contributes to the further formation of low-
534 volatility compounds. Overall, the highest average degree of oxidation ($O/C = 0.77$) is observed in high RH
535 wet aerosol experiment (Expt. 5). The effect of particle water on monoterpene SOA formation warrants
536 further studies.

537

538 3.4 Atmospheric implications

539 There is emerging evidence that monoterpene SOA greatly contribute to atmospheric aerosol in the
540 Southeastern U.S. (Zhang et al., 2018; Xu et al., 2018a). ON is no exception; a substantial fraction of pON
541 is considered to be from monoterpenes (Lee et al., 2016; Huang et al., 2019; Xu et al., 2015a). While C_{10}
542 pON measured in FIGAERO-HR-ToF-I-CIMS is a good tracer of monoterpene derived pON , we show that
543 a fair amount of α -/ β -pinene pON is found as $C_{<10}$ or C_{20} depending on the oxidation condition. This implies
544 that the contribution of monoterpene derived pON could be substantially underestimated when only

545 considering C₁₀ pON. Fraction of pON with different number of carbon reported in this study (Table S2) is
546 a useful parameter to quantitatively determine the contribution of monoterpene derived pON to total pON.

547 Many previous modeling studies using ambient measurement data as constraints report that the
548 lifetime of pON is likely several hours (Pye et al., 2015; Fisher et al., 2016; Lee et al., 2016; Zare et al.,
549 2018). Hydrolysis of pON is used as a dominant loss process with the lifetime of several hours to improve
550 the concentration of modeled OA (Pye et al., 2015) and to improve the concentrations of gaseous and
551 particulate ON (Fisher et al., 2016). If the ambient pON concentration is indeed governed by a loss process
552 with a few hours of lifetime, our results imply that the particle-phase hydrolysis may not be the only
553 dominant loss process because the hydrolysis lifetime reported in this study is significantly shorter. Other
554 potential but less-studied loss mechanisms of ON and pON include deposition (Nguyen et al., 2015),
555 photolysis (Muller et al., 2014), and aqueous photooxidation (Romonosky et al., 2017; Nah et al., 2016).
556 For instance, enhanced photolysis rate is observed for carbonyl nitrate derived from isoprene (Muller et al.,
557 2014), while no similar study is available for monoterpene derived ON in literature. Also, it has been
558 demonstrated that different monoterpene pON can have drastically different photochemical fates (Nah et
559 al., 2016). Taken together, results from this study highlight the importance to investigate other potential
560 loss processes of monoterpene derived pON.

561 Aside from the hydrolysis lifetime, many modeling studies assume F_H as unity (Pye et al., 2015;
562 Fisher et al., 2016; Lee et al., 2016). Even when F_H is considered, the value of F_H used in other studies is
563 still substantially higher than estimated in this study (Browne et al., 2013; Zare et al., 2018). The use of
564 higher F_H would result in overestimating the contribution of hydrolysis as a loss process of pON and NO_x.

565 While hydrolysis is considered as a permanent sink of NO_x, many other loss processes, such as further OH·
566 ~~hydroxyl radical~~-oxidation and photolysis, are considered as a temporary reservoir of NO_x. If the relative
567 importance of pON fates in models was not accurate, the role of ON in NO_x cycling and the formation
568 potential of ozone could have been inaccurately interpreted. Therefore, results from this study regarding
569 the hydrolysis lifetime and F_H serve as experimentally constrained parameters for chemical transport

570 models to accurately evaluate the role of ON in regards to nitrogen budget and the formation of ozone and
571 fine aerosol.

572

573 **Author contribution**

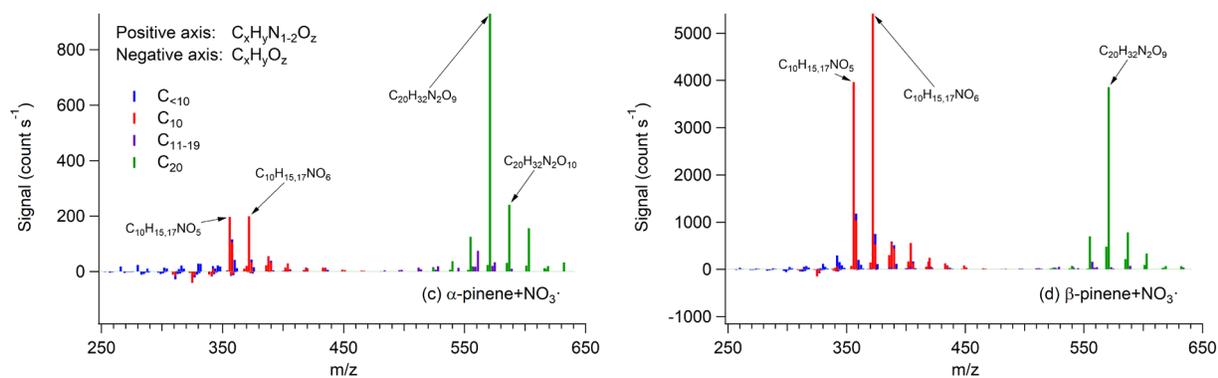
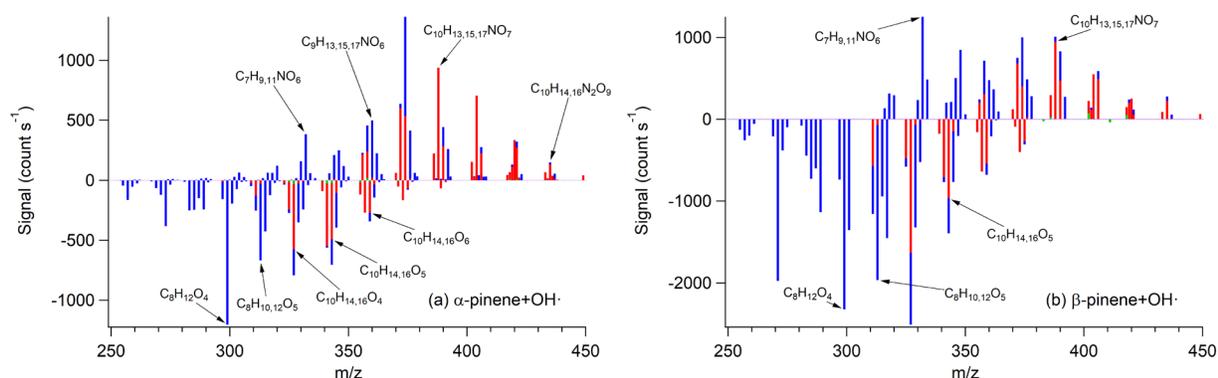
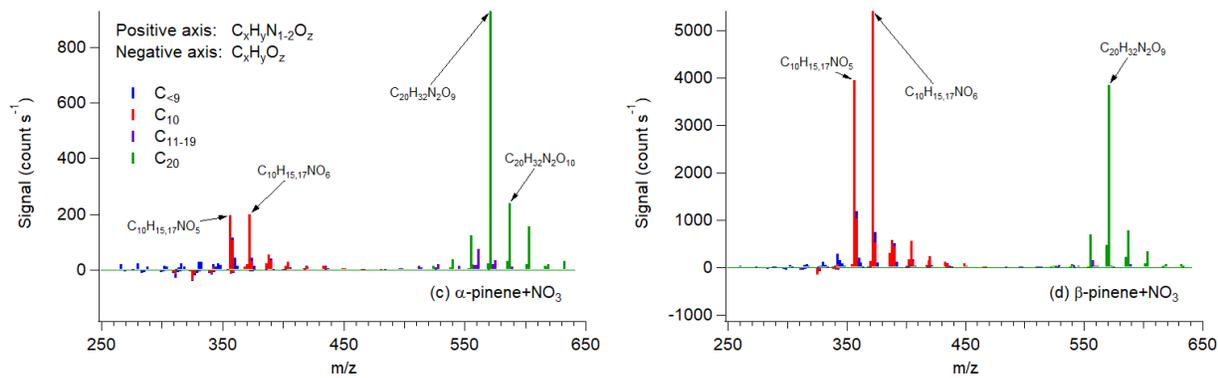
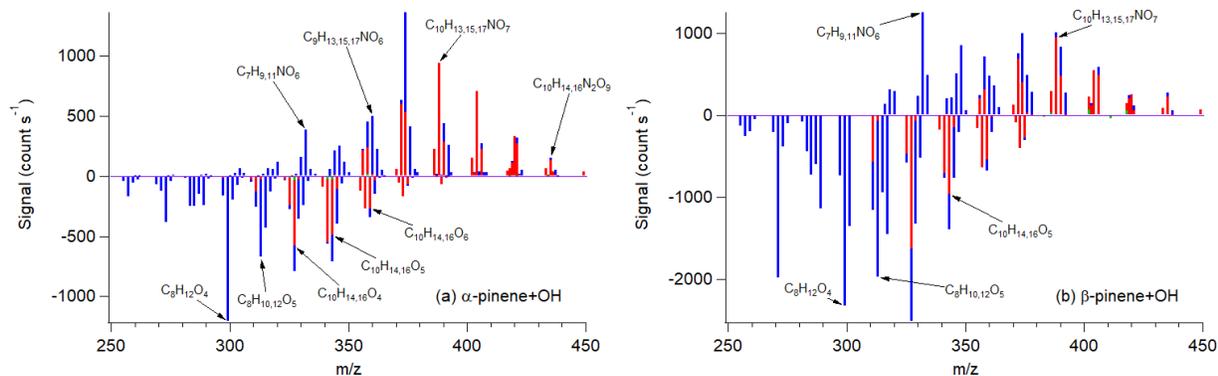
574 M.T. designed and performed the research and analyzed the data with substantial inputs from N.L.N. M.T.
575 and N.L.N. wrote the manuscript.

576

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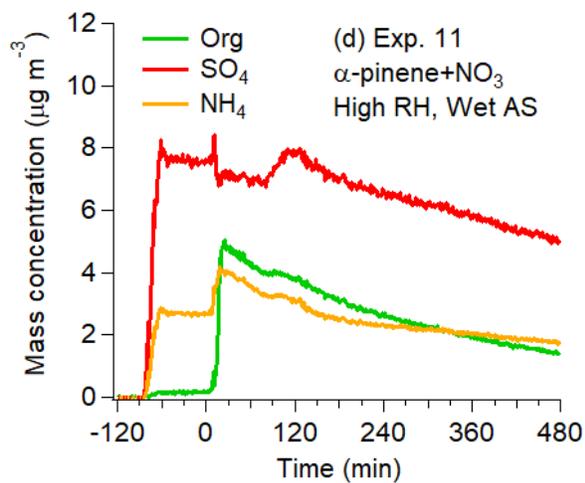
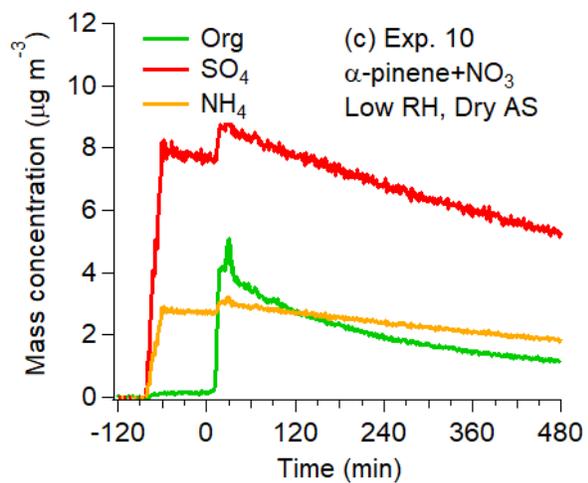
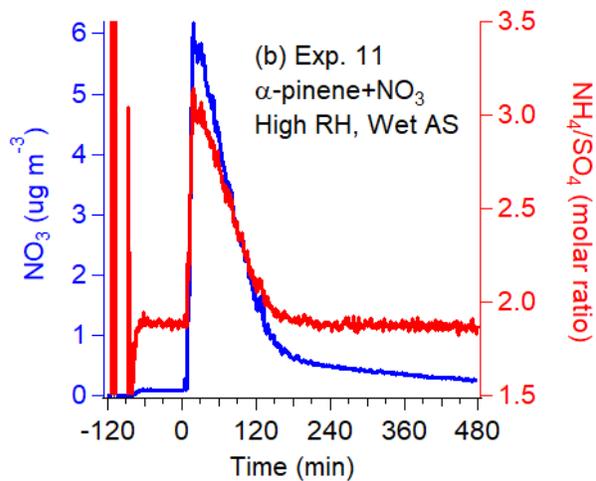
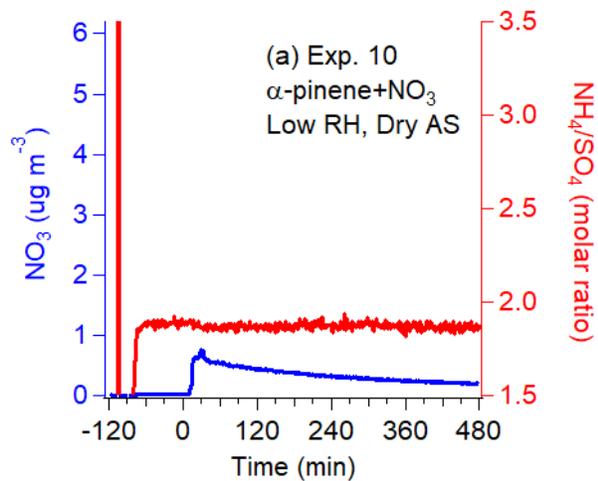
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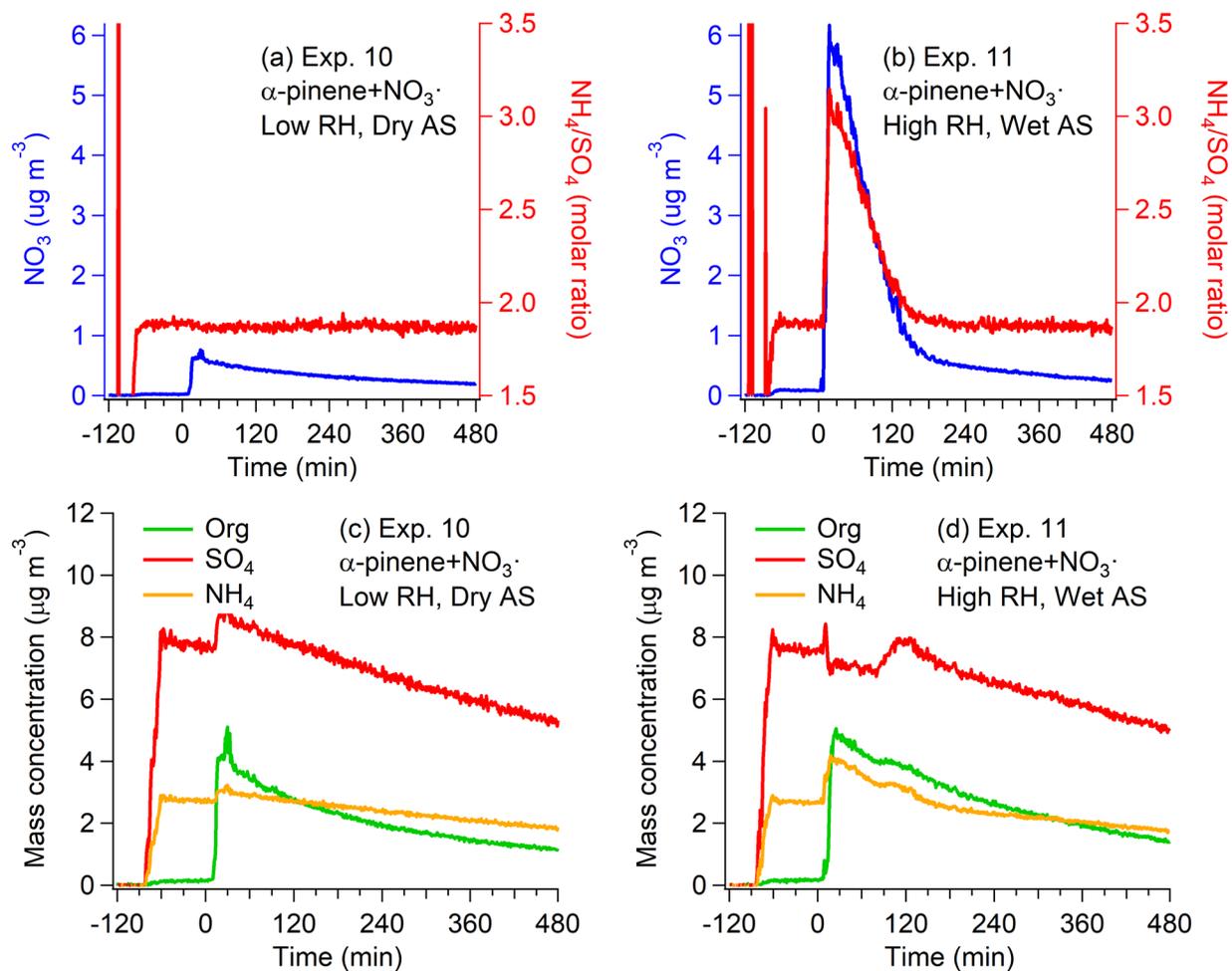
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Figure 1. FIGAERO-HR-ToF-I-CIMS mass spectra of SOA in (a) α -pinene+OH⁻ from Exp. 3, (b) β -

586 pinene+OH₂⁺ from Exp. 6, (c) α-pinene+NO₃⁻ from Exp. 10, and (d) β-pinene+NO₃⁻ from Exp. 14. Top
587 portion of each panel represents C_xH_yN₁₋₂O_z, whereas bottom represents C_xH_yO_z. Bars are colored by the
588 number of carbon as noted in the legend. Prominent masses are labeled with the corresponding chemical
589 formulae without an iodide ion.
590



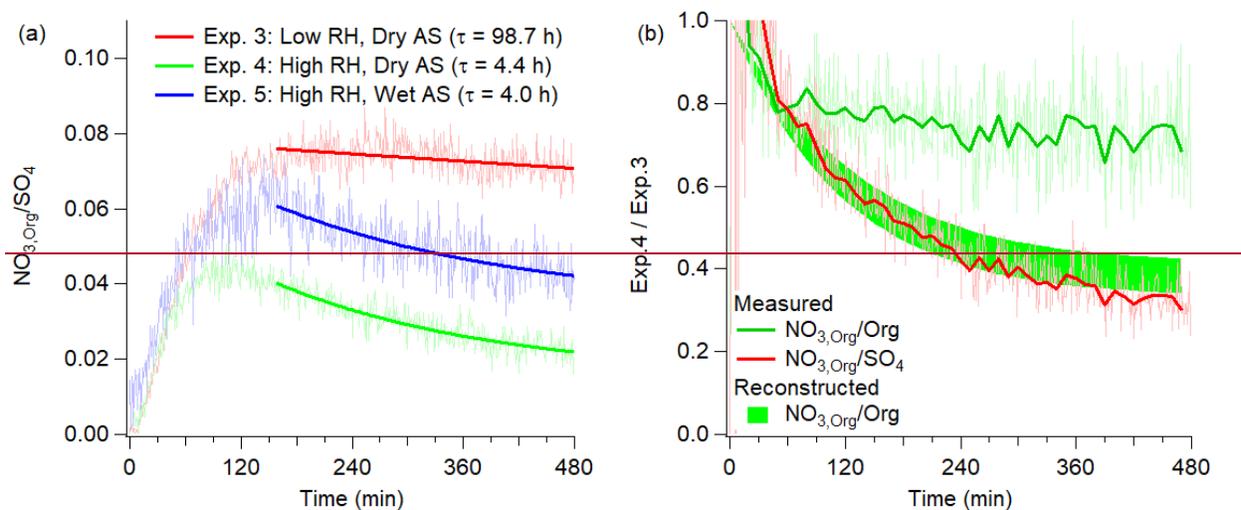


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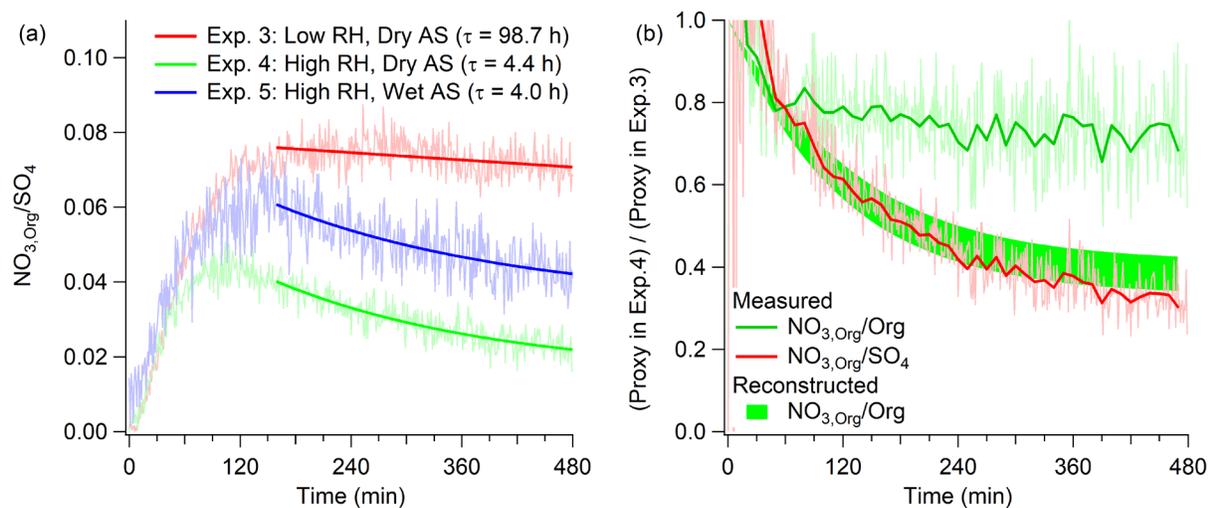
593 Figure 2. (a, b) NO_3 concentration and the molar ratio of NH_4 to SO_4 and (c, d) concentrations of Org, SO_4 ,
 594 and NH_4 measured by HR-ToF-AMS. Data presented in Panels (a) and (c) are from Exp. 10 (low RH, dry
 595 AS), while those in Panels (b) and (d) are from Exp. 11 (high RH, wet AS). These two experiments are
 596 essentially the same except for RH and phase state of seed aerosol.

597

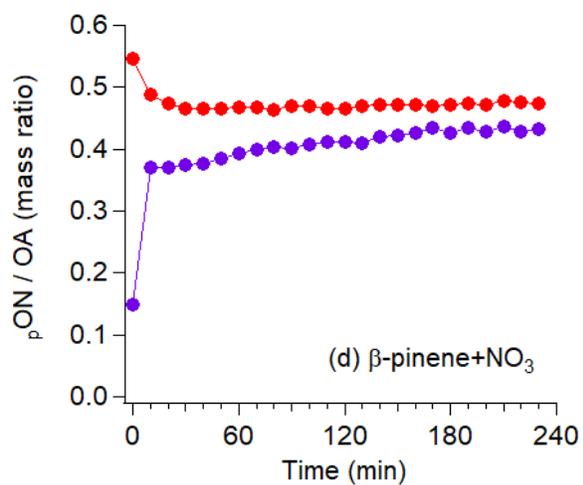
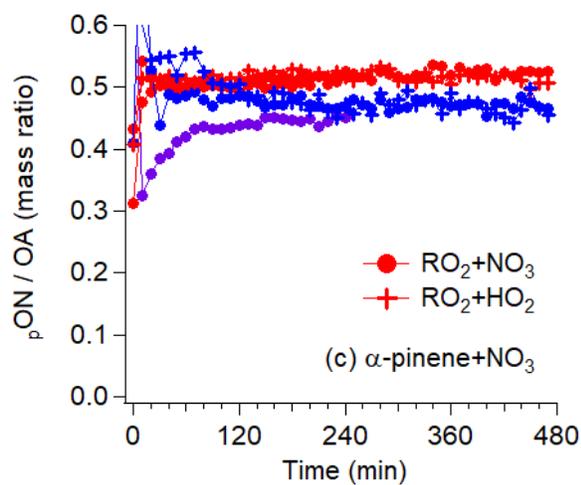
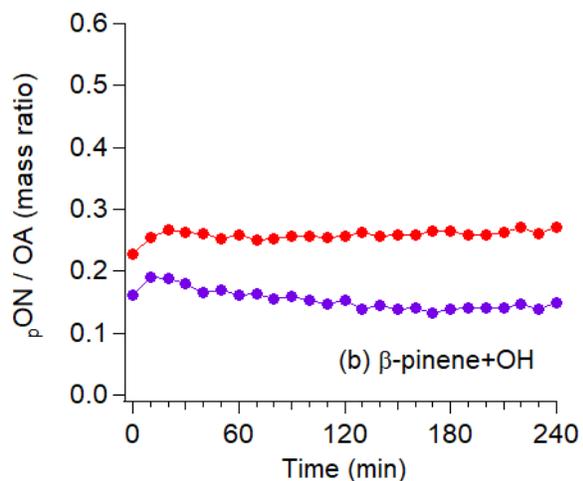
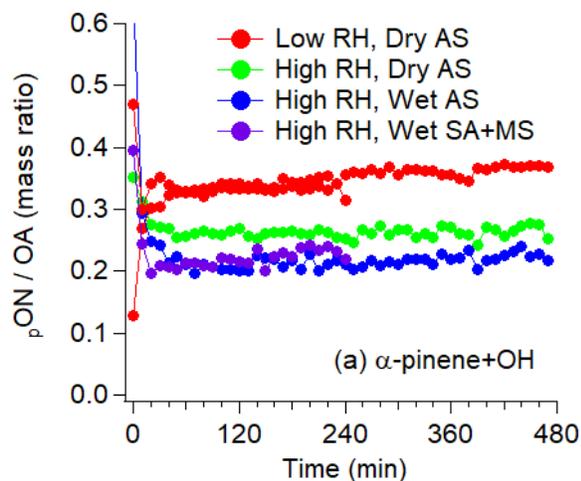
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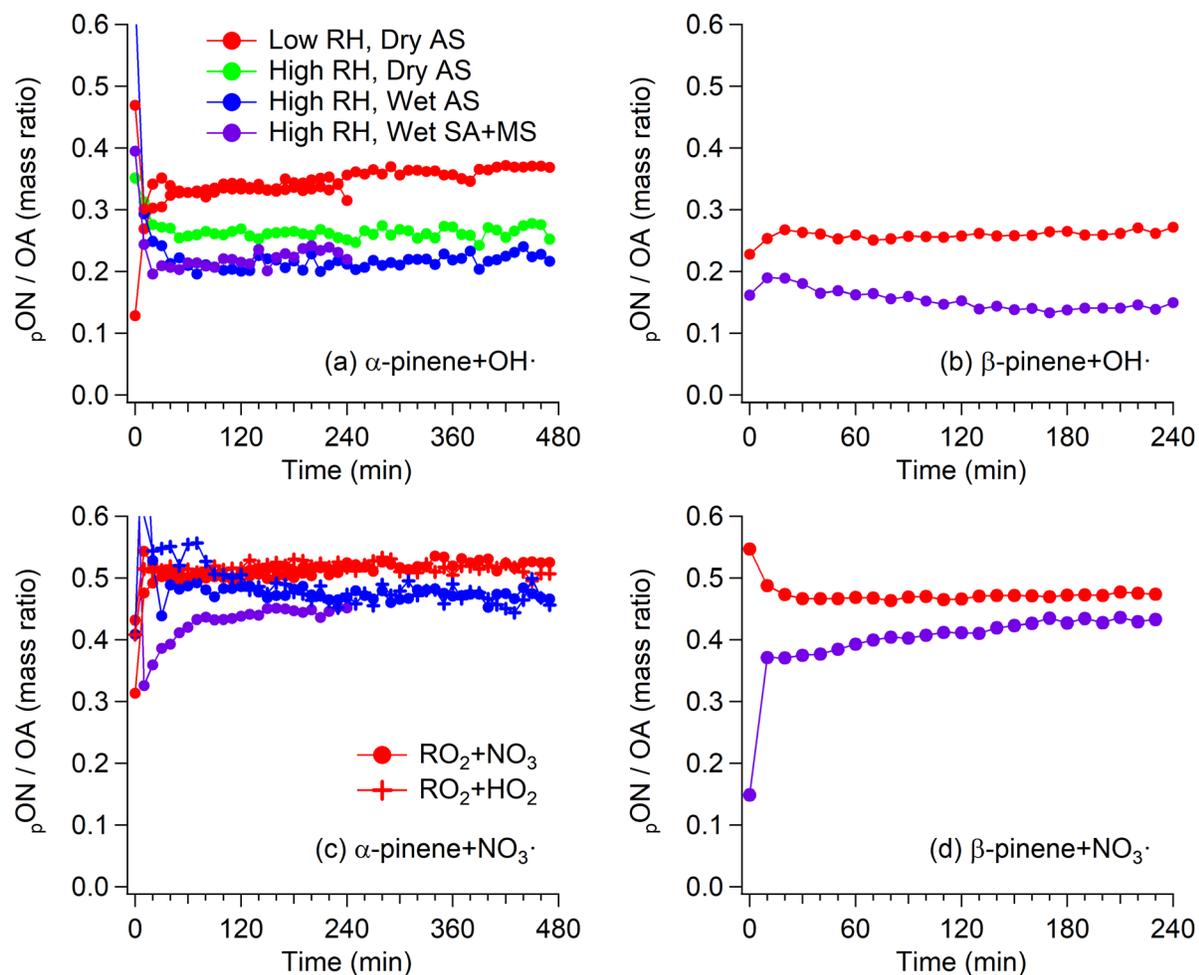


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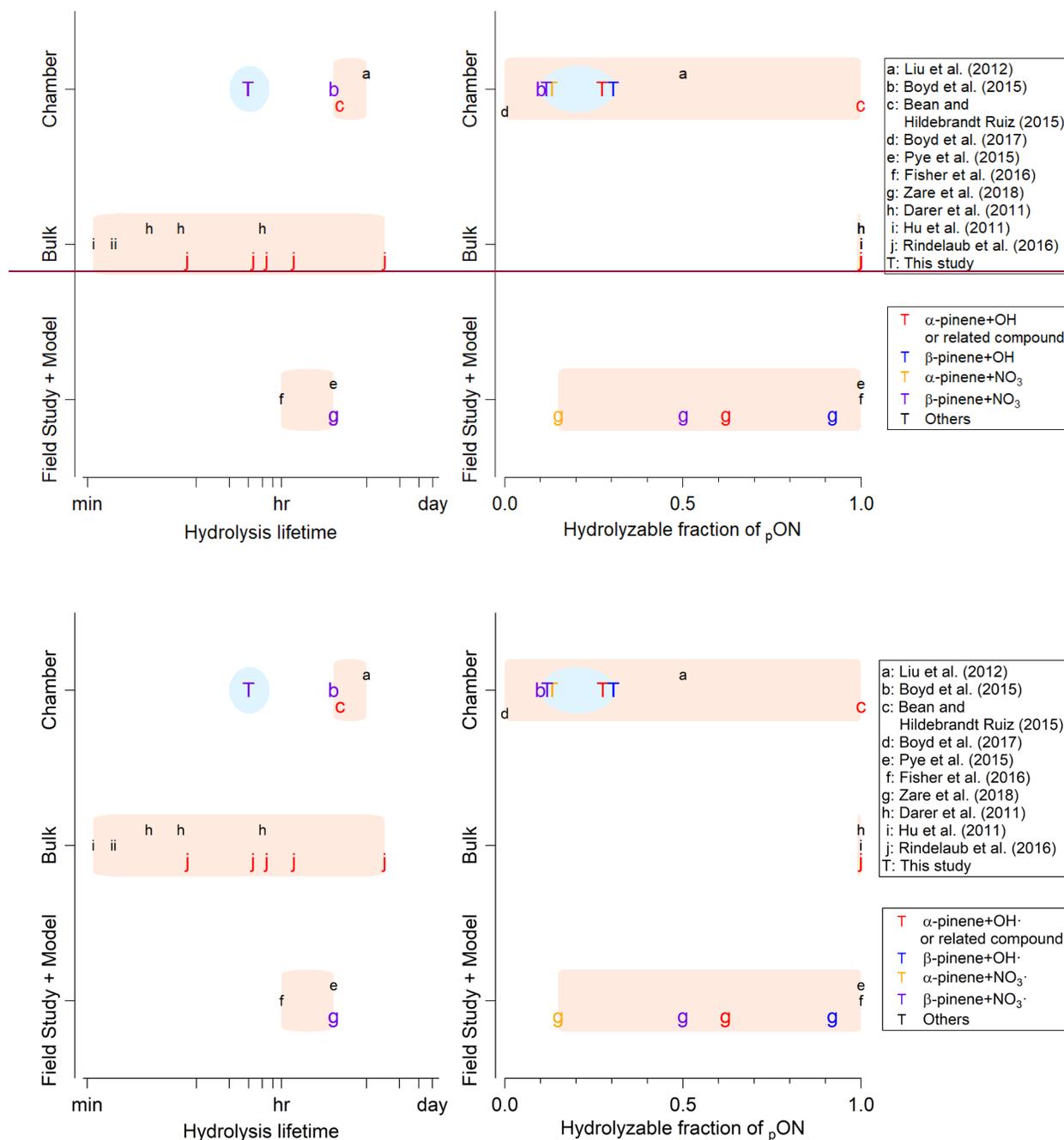
600 Figure 3. (a) Time-series data of $\text{NO}_{3,\text{Org}}/\text{SO}_4$ from Exp. 3-5 (α -pinene+OH \cdot) and the exponential fits with
 601 corresponding characteristic times. (b) $\text{NO}_{3,\text{Org}}/\text{Org}$, $\text{NO}_{3,\text{Org}}/\text{SO}_4$, reconstructed $\text{NO}_{3,\text{Org}}/\text{Org}$ based on the
 602 decay rate of $\text{NO}_{3,\text{Org}}/\text{SO}_4$. Each proxy in Exp. 4 is divided by that in Exp. 3 to determine the relative decay
 603 between high and low RH experiments.





605
 606 Figure 4. Time-series data of p_{ON}/OA in (a) α -pinene+OH₂ from Exp. 1-5, (b) β -pinene+OH₂ from Exp. 6-
 607 7, (c) α -pinene+NO₃ from Exp. 8-13, and (d) β -pinene+NO₃ from Exp. 14-15. Data points are colored by
 608 conditions concerning reactor RH and phase state of seed aerosol. For α -pinene+NO₃, data points are also
 609 shaped differently depending on the expected dominant RO_2 fate.

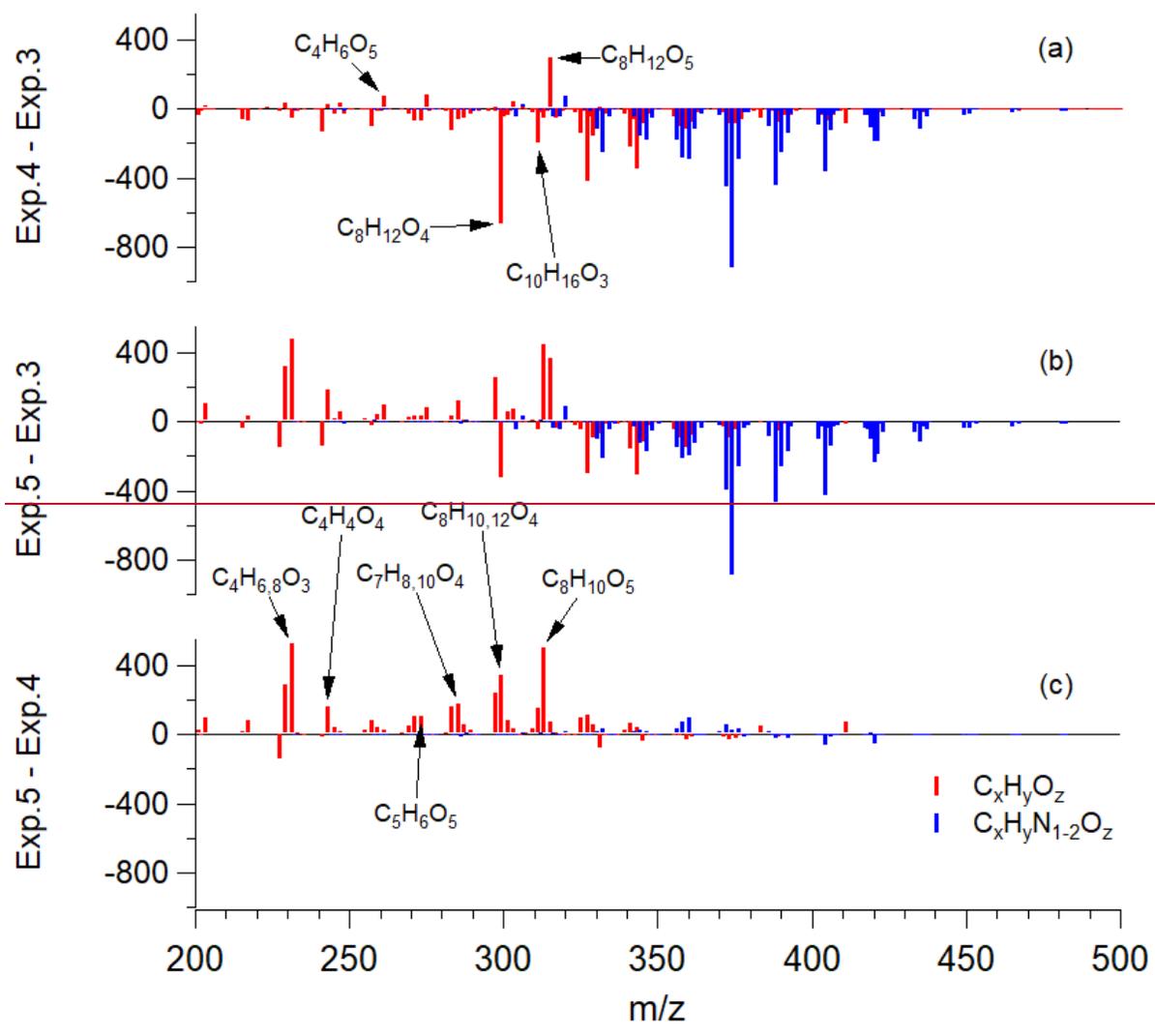
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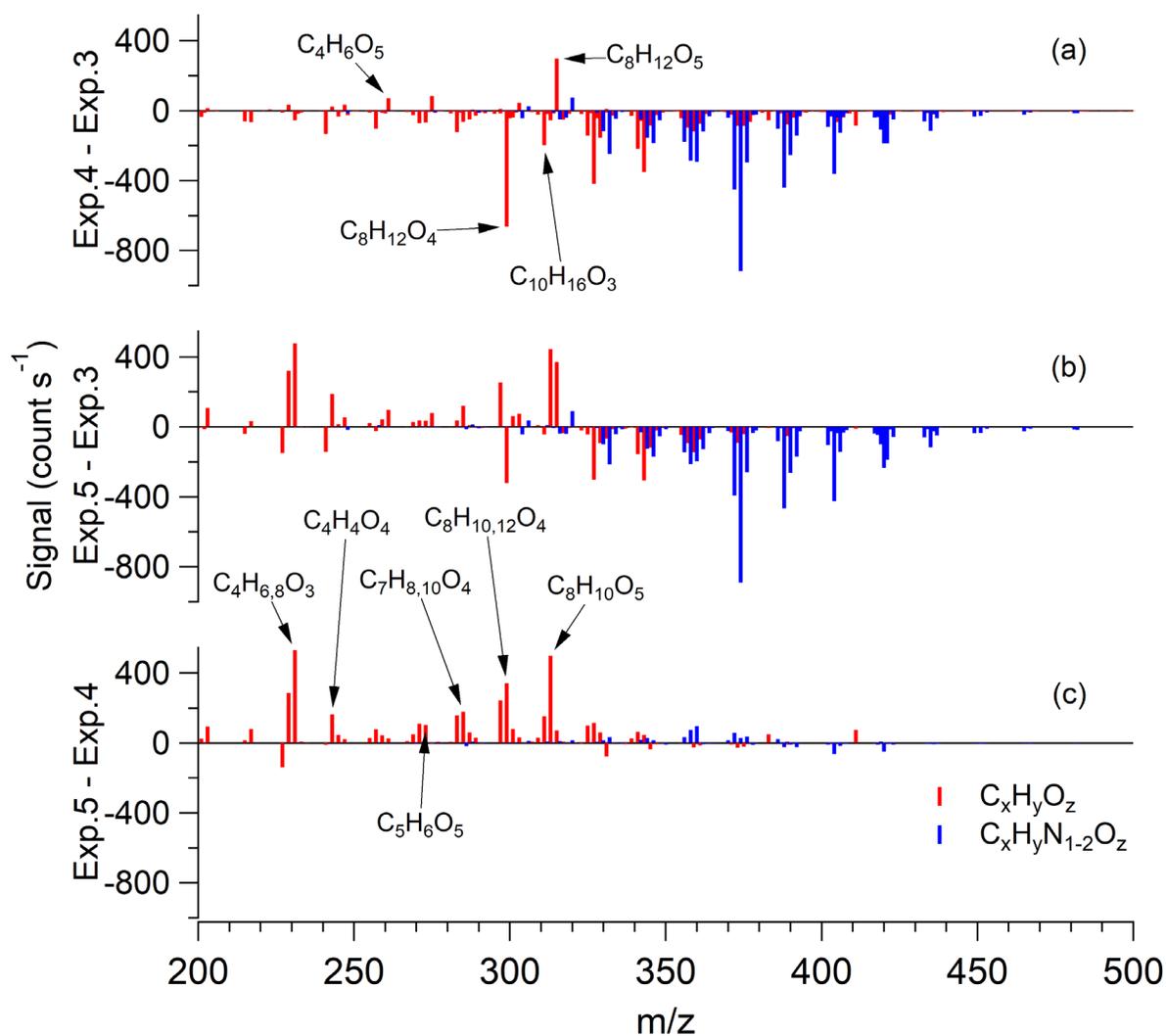
613 Figure 5. Comparison of hydrolysis lifetime of organic nitrate and hydrolyzable fraction of p_{ON} in
 614 literature. “Chamber” refers to laboratory studies of organic nitrate aerosol via chamber experiments,
 615 “Bulk” refers to laboratory studies of organic nitrate compounds using bulk solution measurements, and
 616 “Model” refers to modeling studies using ambient measurement data as constraints. Data points are colored
 617 by the system of VOC and oxidant condition ~~explored in this study~~ and are alphabetized based on the

618 reference. The relevant systems explored in this study are emphasized by enlarging the corresponding font
619 size. Pink shaded regions are ranges reported in literature, while blue shaded regions are ranges reported in
620 this study.

621



622



623

624 Figure 6. FIGAERO-HR-ToF-I-CIMS difference mass spectra of SOA in α -pinene+OH. (a) Exp. 4 (high

625 RH, dry AS) minus Exp. 3 (low RH, dry AS). (b) Exp. 5 (high RH, wet AS) minus Exp. 3 (low RH, dry

626 AS). (c) Exp. 5 (high RH, wet AS) minus Exp. 4 (high RH, dry AS). Bars are colored by the difference in

627 chemical composition (i.e., red for $C_xH_yO_z$ and blue for $C_xH_yN_{1-2}O_z$). Prominent masses are labeled with

628 the corresponding chemical formulae without an iodide ion.

629

630 Table 1. Summary of experimental conditions considered in this study.

Exp.	Precursor VOC	Oxidant precursor	Reactor RH	Seed
1	α -pinene (25 ppb)	H ₂ O ₂ (2 ppm), NO (6 ppb min ⁻¹)	~51-3 %	Effloresced AS ^b
2	α -pinene (25 ppb)	H ₂ O ₂ (2 ppm), NO (6 ppb min ⁻¹)	~50-70 48- <u>65</u> % ^a	Deliquesced SA+MS ^c
3	α -pinene (25 ppb)	H ₂ O ₂ (2 ppm), NO (2 ppb min ⁻¹)	~52-6 %	Effloresced AS
4	α -pinene (25 ppb)	H ₂ O ₂ (2 ppm), NO (2 ppb min ⁻¹)	~50-70 53- <u>66</u> % ^a	Effloresced AS
5	α -pinene (25 ppb)	H ₂ O ₂ (2 ppm), NO (2 ppb min ⁻¹)	~50-70 57- <u>72</u> % ^a	Deliquesced AS
6	β -pinene (25 ppb)	H ₂ O ₂ (2 ppm), NO (6 ppb min ⁻¹)	~51-3 %	Effloresced AS
7	β -pinene (25 ppb)	H ₂ O ₂ (2 ppm), NO (6 ppb min ⁻¹)	~50-70 53- <u>70</u> % ^a	Deliquesced SA+MS
8	α -pinene (12 ppb)	N ₂ O ₅ (80 ppb)	~52-3 %	Effloresced AS
9	α -pinene (12 ppb)	N ₂ O ₅ (80 ppb)	~60-70 67- <u>71</u> %	Deliquesced SA+MS
10	α -pinene (12 ppb)	N ₂ O ₅ (80 ppb)	~51-6 %	Effloresced AS
11	α -pinene (12 ppb)	N ₂ O ₅ (80 ppb)	~60-70 69- <u>74</u> %	Deliquesced AS
12	α -pinene (12 ppb)	N ₂ O ₅ (80 ppb), HCHO (25 ppm)	~53-8 %	Effloresced AS
13	α -pinene (12 ppb)	N ₂ O ₅ (80 ppb), HCHO (25 ppm)	~60-70 67- <u>71</u> %	Deliquesced AS
14	β -pinene (12 ppb)	N ₂ O ₅ (80 ppb)	~52-5 %	Effloresced AS
15	β -pinene (12 ppb)	N ₂ O ₅ (80 ppb)	~60-70 56- <u>72</u> %	Deliquesced SA+MS

631 ^aThe target for the initial reactor RH is ~70 %. However, the irradiation of UV light increases the reactor
 632 temperature by several degree Celsius and hence decreases RH.

633 ^bAmmonium sulfate

634 ^cSulfuric acid and magnesium sulfate

635

636 Table 2. Hydrolysis lifetime and corresponding fraction of hydrolyzable pON.

System	Hydrolysis lifetime	Hydrolyzable fraction (F _H) ^a
α-pinene+OH ₂	<30 min	23–32 %
β-pinene+OH ₂	<30 min	27–34 %
α-pinene+NO ₃	<30 min	9–17 %
β-pinene+NO ₃	<30 min	9–15 %

637 ^aLower or higher limit of hydrolyzable pON fraction is based on the assumptions that organic moiety of
 638 hydrolysis products remains in aerosol or partitions back to the gas phase.

639

640 **References**

- 641 Aljawhary, D., Lee, A. K. Y., and Abbatt, J. P. D.: High-resolution chemical ionization mass spectrometry
642 (ToF-CIMS): application to study SOA composition and processing, *Atmos Meas Tech*, 6, 3211-3224,
643 10.5194/amt-6-3211-2013, 2013.
- 644 Aljawhary, D., Zhao, R., Lee, A. K. Y., Wang, C., and Abbatt, J. P. D.: Kinetics, Mechanism, and Secondary
645 Organic Aerosol Yield of Aqueous Phase Photo-oxidation of alpha-Pinene Oxidation Products, *J Phys*
646 *Chem A*, 120, 1395-1407, 10.1021/acs.jpca.5b06237, 2016.
- 647 Aschmann, S. M., Atkinson, R., and Arey, J.: Products of reaction of OH radicals with alpha-pinene, *J*
648 *Geophys Res-Atmos*, 107, Artn 4191, 10.1029/2001jd001098, 2002.
- 649 Atkinson, R., and Arey, J.: Atmospheric degradation of volatile organic compounds, *Chem Rev*, 103, 4605-
650 4638, 10.1021/cr0206420, 2003.
- 651 Bean, J. K., and Hildebrandt Ruiz, L.: Gas-particle partitioning and hydrolysis of organic nitrates formed
652 from the oxidation of alpha-pinene in environmental chamber experiments, *Atmos Chem Phys*, 16, 2175-
653 2184, 10.5194/acp-16-2175-2016, 2016.
- 654 Berndt, T., and Boge, O.: Products and mechanism of the gas-phase reaction of NO₃ radicals with alpha-
655 pinene, *J Chem Soc Faraday T*, 93, 3021-3027, DOI 10.1039/a702364b, 1997.
- 656 Berndt, T., Richters, S., Jokinen, T., Hyttinen, N., Kurten, T., Otkjaer, R. V., Kjaergaard, H. G., Stratmann,
657 F., Herrmann, H., Sipila, M., Kulmala, M., and Ehn, M.: Hydroxyl radical-induced formation of highly
658 oxidized organic compounds, *Nat Commun*, 7, 10.1038/ncomms13677, 2016.
- 659 Bertram, T. H., Kimmel, J. R., Crisp, T. A., Ryder, O. S., Yatavelli, R. L. N., Thornton, J. A., Cubison, M.
660 J., Gonin, M., and Worsnop, D. R.: A field-deployable, chemical ionization time-of-flight mass
661 spectrometer, *Atmos Meas Tech*, 4, 1471-1479, 10.5194/amt-4-1471-2011, 2011.
- 662 Boschan, R., Merrow, R. T., and Vandolah, R. W.: The Chemistry of Nitrate Esters, *Chem Rev*, 55, 485-
663 510, DOI 10.1021/cr50003a001, 1955.
- 664 Boyd, C. M., Sanchez, J., Xu, L., Eugene, A. J., Nah, T., Tuet, W. Y., Guzman, M. I., and Ng, N. L.:
665 Secondary organic aerosol formation from the beta-pinene+NO₃ system: effect of humidity and peroxy
666 radical fate, *Atmos Chem Phys*, 15, 7497-7522, 10.5194/acp-15-7497-2015, 2015.
- 667 Boyd, C. M., Nah, T., Xu, L., Berkemeier, T., and Ng, N. L.: Secondary Organic Aerosol (SOA) from
668 Nitrate Radical Oxidation of Monoterpenes: Effects of Temperature, Dilution, and Humidity on Aerosol
669 Formation, Mixing, and Evaporation, *Environ Sci Technol*, 51, 7831-7841, 10.1021/acs.est.7b01460, 2017.
- 670 Browne, E. C., Min, K. E., Wooldridge, P. J., Apel, E., Blake, D. R., Brune, W. H., Cantrell, C. A., Cubison,
671 M. J., Diskin, G. S., Jimenez, J. L., Weinheimer, A. J., Wennberg, P. O., Wisthaler, A., and Cohen, R. C.:
672 Observations of total RONO₂ over the boreal forest: NO_x sinks and HNO₃ sources, *Atmos Chem Phys*,
673 13, 4543-4562, 10.5194/acp-13-4543-2013, 2013.
- 674 Bruns, E. A., Perraud, V., Zelenyuk, A., Ezell, M. J., Johnson, S. N., Yu, Y., Imre, D., Finlayson-Pitts, B.
675 J., and Alexander, M. L.: Comparison of FTIR and Particle Mass Spectrometry for the Measurement of
676 Particulate Organic Nitrates, *Environ Sci Technol*, 44, 1056-1061, 10.1021/es9029864, 2010.

677 Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P., Hildebrandt Ruiz, L.,
678 Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue, N. M., Jayne, J. T., and Worsnop, D.
679 R.: Elemental ratio measurements of organic compounds using aerosol mass spectrometry: characterization,
680 improved calibration, and implications, *Atmos Chem Phys*, 15, 253-272, 10.5194/acp-15-253-2015, 2015.

681 Carlton, A. G., Pinder, R. W., Bhave, P. V., and Pouliot, G. A.: To What Extent Can Biogenic SOA be
682 Controlled?, *Environ Sci Technol*, 44, 3376-3380, 10.1021/es903506b, 2010.

683 Cerully, K. M., Bougiatioti, A., Hite, J. R., Guo, H., Xu, L., Ng, N. L., Weber, R., and Nenes, A.: On the
684 link between hygroscopicity, volatility, and oxidation state of ambient and water-soluble aerosols in the
685 southeastern United States, *Atmos Chem Phys*, 15, 8679-8694, 10.5194/acp-15-8679-2015, 2015.

686 Clafin, M. S., and Ziemann, P. J.: Identification and Quantitation of Aerosol Products of the Reaction of
687 beta-Pinene with NO₃ Radicals and Implications for Gas- and Particle-Phase Reaction Mechanisms, *J Phys
688 Chem A*, 122, 3640-3652, 10.1021/acs.jpca.8b00692, 2018.

689 Crouse, J. D., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., and Wennberg, P. O.: Autoxidation of
690 Organic Compounds in the Atmosphere, *J Phys Chem Lett*, 4, 3513-3520, 10.1021/jz4019207, 2013.

691 Darer, A. I., Cole-Filipiak, N. C., O'Connor, A. E., and Elrod, M. J.: Formation and Stability of
692 Atmospherically Relevant Isoprene-Derived Organosulfates and Organonitrates, *Environ Sci Technol*, 45,
693 1895-1902, 10.1021/es103797z, 2011.

694 Day, D. A., Liu, S., Russell, L. M., and Ziemann, P. J.: Organonitrate group concentrations in submicron
695 particles with high nitrate and organic fractions in coastal southern California, *Atmos Environ*, 44, 1970-
696 1979, 10.1016/j.atmosenv.2010.02.045, 2010.

697 DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C., Gonin, M., Fuhrer,
698 K., Horvath, T., Docherty, K. S., Worsnop, D. R., and Jimenez, J. L.: Field-deployable, high-resolution,
699 time-of-flight aerosol mass spectrometer, *Anal Chem*, 78, 8281-8289, 10.1021/ac061249n, 2006.

700 Eddingsaas, N. C., Loza, C. L., Yee, L. D., Chan, M., Schilling, K. A., Chhabra, P. S., Seinfeld, J. H., and
701 Wennberg, P. O.: alpha-pinene photooxidation under controlled chemical conditions - Part 2: SOA yield
702 and composition in low- and high-NO_x environments, *Atmos Chem Phys*, 12, 7413-7427, 10.5194/acp-12-
703 7413-2012, 2012.

704 Ehn, M., Thornton, J. A., Kleist, E., Sipila, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F.,
705 Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I. H., Rissanen, M., Jokinen, T., Schobesberger,
706 S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurten, T., Nielsen, L. B., Jorgensen, S., Kjaergaard, H.
707 G., Canagaratna, M., Dal Maso, M., Berndt, T., Petaja, T., Wahner, A., Kerminen, V. M., Kulmala, M.,
708 Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol,
709 *Nature*, 506, 476+, 10.1038/nature13032, 2014.

710 Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in cloud droplets and
711 aqueous particles (aqSOA): a review of laboratory, field and model studies, *Atmos Chem Phys*, 11, 11069-
712 11102, 10.5194/acp-11-11069-2011, 2011.

713 Farmer, D. K., Matsunaga, A., Docherty, K. S., Surratt, J. D., Seinfeld, J. H., Ziemann, P. J., and Jimenez,
714 J. L.: Response of an aerosol mass spectrometer to organonitrates and organosulfates and implications for
715 atmospheric chemistry, *P Natl Acad Sci USA*, 107, 6670-6675, 10.1073/pnas.0912340107, 2010.

716 Faxon, C., Hammes, J., Le Breton, M., Pathak, R. K., and Hallquist, M.: Characterization of organic nitrate
717 constituents of secondary organic aerosol (SOA) from nitrate-radical-initiated oxidation of limonene using
718 high-resolution chemical ionization mass spectrometry, *Atmos Chem Phys*, 18, 5467-5481, 10.5194/acp-
719 18-5467-2018, 2018.

720 Fisher, J. A., Jacob, D. J., Travis, K. R., Kim, P. S., Marais, E. A., Miller, C. C., Yu, K. R., Zhu, L.,
721 Yantosca, R. M., Sulprizio, M. P., Mao, J. Q., Wennberg, P. O., Crouse, J. D., Teng, A. P., Nguyen, T. B.,
722 St Clair, J. M., Cohen, R. C., Romer, P., Nault, B. A., Wooldridge, P. J., Jimenez, J. L., Campuzano-Jost,
723 P., Day, D. A., Hu, W. W., Shepson, P. B., Xiong, F. L. Z., Blake, D. R., Goldstein, A. H., Misztal, P. K.,
724 Hanisco, T. F., Wolfe, G. M., Ryerson, T. B., Wisthaler, A., and Mikoviny, T.: Organic nitrate chemistry
725 and its implications for nitrogen budgets in an isoprene- and monoterpene-rich atmosphere: constraints
726 from aircraft (SEAC(4)RS) and ground-based (SOAS) observations in the Southeast US, *Atmos Chem*
727 *Phys*, 16, 5969-5991, 10.5194/acp-16-5969-2016, 2016.

728 Fry, J. L., Kiendler-Scharr, A., Rollins, A. W., Wooldridge, P. J., Brown, S. S., Fuchs, H., Dube, W.,
729 Mensah, A., dal Maso, M., Tillmann, R., Dorn, H. P., Brauers, T., and Cohen, R. C.: Organic nitrate and
730 secondary organic aerosol yield from NO₃ oxidation of beta-pinene evaluated using a gas-phase
731 kinetics/aerosol partitioning model, *Atmos Chem Phys*, 9, 1431-1449, DOI 10.5194/acp-9-1431-2009,
732 2009.

733 Fry, J. L., Draper, D. C., Zarzana, K. J., Campuzano-Jost, P., Day, D. A., Jimenez, J. L., Brown, S. S.,
734 Cohen, R. C., Kaser, L., Hansel, A., Cappellin, L., Karl, T., Roux, A. H., Turnipseed, A., Cantrell, C., Lefer,
735 B. L., and Grossberg, N.: Observations of gas- and aerosol-phase organic nitrates at BEACHON-RoMBAS
736 2011, *Atmos Chem Phys*, 13, 8585-8605, 10.5194/acp-13-8585-2013, 2013.

737 Fry, J. L., Draper, D. C., Barsanti, K. C., Smith, J. N., Ortega, J., Winkle, P. M., Lawler, M. J., Brown, S.
738 S., Edwards, P. M., Cohen, R. C., and Lee, L.: Secondary Organic Aerosol Formation and Organic Nitrate
739 Yield from NO₃ Oxidation of Biogenic Hydrocarbons, *Environ Sci Technol*, 48, 11944-11953,
740 10.1021/es502204x, 2014.

741 Fry, J. L., Brown, S. S., Middlebrook, A. M., Edwards, P. M., Campuzano-Jost, P., Day, D. A., Jimenez, J.
742 L., Allen, H. M., Ryerson, T. B., Pollack, I., Graus, M., Warneke, C., de Gouw, J. A., Brock, C. A., Gilman,
743 J., Lerner, B. M., Dube, W. P., Liao, J., and Welti, A.: Secondary organic aerosol (SOA) yields from NO₃
744 radical + isoprene based on nighttime aircraft power plant plume transects, *Atmos Chem Phys*, 18, 11663-
745 11682, 10.5194/acp-18-11663-2018, 2018.

746 Gao, S., Ng, N. L., Keywood, M., Varutbangkul, V., Bahreini, R., Nenes, A., He, J. W., Yoo, K. Y.,
747 Beauchamp, J. L., Hodyss, R. P., Flagan, R. C., and Seinfeld, J. H.: Particle phase acidity and oligomer
748 formation in secondary organic aerosol, *Environ Sci Technol*, 38, 6582-6589, 10.1021/es049125k, 2004.

749 Goldstein, A. H., and Galbally, I. E.: Known and unexplored organic constituents in the earth's atmosphere,
750 *Environ Sci Technol*, 41, 1514-1521, DOI 10.1021/es072476p, 2007.

751 Griffin, R. J., Cocker, D. R., Flagan, R. C., and Seinfeld, J. H.: Organic aerosol formation from the oxidation
752 of biogenic hydrocarbons, *J Geophys Res-Atmos*, 104, 3555-3567, Doi 10.1029/1998jd100049, 1999.

753 Grzanic, G., Bartels-Rausch, T., Berkemeier, T., Turler, A., and Ammann, M.: Viscosity controls humidity
754 dependence of N₂O₅ uptake to citric acid aerosol, *Atmos Chem Phys*, 15, 13615-13625, 10.5194/acp-15-
755 13615-2015, 2015.

756 Guenther, A. B., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T., Emmons, L. K., and Wang, X.:
757 The Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1): an extended and
758 updated framework for modeling biogenic emissions, *Geosci Model Dev*, 5, 1471-1492, 10.5194/gmd-5-
759 1471-2012, 2012.

760 Hallquist, M., Wangberg, I., Ljungstrom, E., Barnes, I., and Becker, K. H.: Aerosol and product yields from
761 NO₃ radical-initiated oxidation of selected monoterpenes, *Environ Sci Technol*, 33, 553-559, DOI
762 10.1021/es980292s, 1999.

763 Hoyle, C. R., Boy, M., Donahue, N. M., Fry, J. L., Glasius, M., Guenther, A., Hallar, A. G., Hartz, K. H.,
764 Petters, M. D., Petaja, T., Rosenoern, T., and Sullivan, A. P.: A review of the anthropogenic influence on
765 biogenic secondary organic aerosol, *Atmos Chem Phys*, 11, 321-343, 10.5194/acp-11-321-2011, 2011.

766 Hu, K. S., Darer, A. I., and Elrod, M. J.: Thermodynamics and kinetics of the hydrolysis of atmospherically
767 relevant organonitrates and organosulfates, *Atmos Chem Phys*, 11, 8307-8320, 10.5194/acp-11-8307-2011,
768 2011.

769 Huang, W., Saathoff, H., Shen, X. L., Ramisetty, R., Leisner, T., and Mohr, C.: Chemical Characterization
770 of Highly Functionalized Organonitrates Contributing to Night-Time Organic Aerosol Mass Loadings and
771 Particle Growth, *Environ Sci Technol*, 53, 1165-1174, 10.1021/acs.est.8b05826, 2019.

772 Huang, Y. L., Zhao, R., Charan, S. M., Kenseth, C. M., Zhang, X., and Seinfeld, J. H.: Unified Theory of
773 Vapor-Wall Mass Transport in Teflon-Walled Environmental Chambers, *Environ Sci Technol*, 52, 2134-
774 2142, 10.1021/acs.est.7b05575, 2018.

775 Jacobs, M. I., Burke, W. J., and Elrod, M. J.: Kinetics of the reactions of isoprene-derived hydroxynitrates:
776 gas phase epoxide formation and solution phase hydrolysis, *Atmos Chem Phys*, 14, 8933-8946,
777 10.5194/acp-14-8933-2014, 2014.

778 Jenkin, M. E., Saunders, S. M., and Pilling, M. J.: The tropospheric degradation of volatile organic
779 compounds: A protocol for mechanism development, *Atmos Environ*, 31, 81-104, Doi 10.1016/S1352-
780 2310(96)00105-7, 1997.

781 Jokinen, T., Berndt, T., Makkonen, R., Kerminen, V. M., Junninen, H., Paasonen, P., Stratmann, F.,
782 Herrmann, H., Guenther, A. B., Worsnop, D. R., Kulmala, M., Ehn, M., and Sipila, M.: Production of
783 extremely low volatile organic compounds from biogenic emissions: Measured yields and atmospheric
784 implications, *P Natl Acad Sci USA*, 112, 7123-7128, 10.1073/pnas.1423977112, 2015.

785 Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., Van Dingenen,
786 R., Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski, Y., Fuzzi, S., Horth, J.,
787 Moortgat, G. K., Winterhalter, R., Myhre, C. E. L., Tsigaridis, K., Vignati, E., Stephanou, E. G., and
788 Wilson, J.: Organic aerosol and global climate modelling: a review, *Atmos Chem Phys*, 5, 1053-1123, DOI
789 10.5194/acp-5-1053-2005, 2005.

790 Kebabian, P. L., Herndon, S. C., and Freedman, A.: Detection of nitrogen dioxide by cavity attenuated
791 phase shift spectroscopy, *Anal Chem*, 77, 724-728, 10.1021/ac048715y, 2005.

792 Kiendler-Scharr, A., Mensah, A. A., Friese, E., Topping, D., Nemitz, E., Prevot, A. S. H., Aijala, M., Allan,
793 J., Canonaco, F., Canagaratna, M., Carbone, S., Crippa, M., Dall'Osto, M., Day, D. A., De Carlo, P., Di
794 Marco, C. F., Elbern, H., Eriksson, A., Freney, E., Hao, L., Herrmann, H., Hildebrandt, L., Hillamo, R.,
795 Jimenez, J. L., Laaksonen, A., McFiggans, G., Mohr, C., O'Dowd, C., Otjes, R., Ovadnevaite, J., Pandis, S.

796 N., Poulain, L., Schlag, P., Sellegri, K., Swietlicki, E., Tiitta, P., Vermeulen, A., Wahner, A., Worsnop, D.,
797 and Wu, H. C.: Ubiquity of organic nitrates from nighttime chemistry in the European submicron aerosol,
798 *Geophys Res Lett*, 43, 7735-7744, 10.1002/2016gl069239, 2016.

799 Krechmer, J. E., Pagonis, D., Ziemann, P. J., and Jimenez, J. L.: Quantification of Gas-Wall Partitioning in
800 Teflon Environmental Chambers Using Rapid Bursts of Low-Volatility Oxidized Species Generated in
801 Situ, *Environ Sci Technol*, 50, 5757-5765, 10.1021/acs.est.6b00606, 2016.

802 Kroll, J. H., and Seinfeld, J. H.: Chemistry of secondary organic aerosol: Formation and evolution of low-
803 volatility organics in the atmosphere, *Atmos Environ*, 42, 3593-3624, 10.1016/j.atmosenv.2008.01.003,
804 2008.

805 Kurten, T., Moller, K. H., Nguyen, T. B., Schwantes, R. H., Misztal, P. K., Su, L. P., Wennberg, P. O., Fry,
806 J. L., and Kjaergaard, H. G.: Alkoxy Radical Bond Scissions Explain the Anomalously Low Secondary
807 Organic Aerosol and Organonitrate Yields From alpha-Pinene + NO₃, *J Phys Chem Lett*, 8, 2826-2834,
808 10.1021/acs.jpcclett.7b01038, 2017.

809 La, Y. S., Camredon, M., Ziemann, P. J., Valorso, R., Matsunaga, A., Lannuque, V., Lee-Taylor, J., Hodzic,
810 A., Madronich, S., and Aumont, B.: Impact of chamber wall loss of gaseous organic compounds on
811 secondary organic aerosol formation: explicit modeling of SOA formation from alkane and alkene
812 oxidation, *Atmos Chem Phys*, 16, 1417-1431, 10.5194/acp-16-1417-2016, 2016.

813 Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Kurten, T., Worsnop, D. R., and Thornton, J. A.: An Iodide-
814 Adduct High-Resolution Time-of-Flight Chemical-Ionization Mass Spectrometer: Application to
815 Atmospheric Inorganic and Organic Compounds, *Environ Sci Technol*, 48, 6309-6317, 10.1021/es500362a,
816 2014.

817 Lee, B. H., Mohr, C., Lopez-Hilfiker, F. D., Lutz, A., Hallquist, M., Lee, L., Romer, P., Cohen, R. C., Iyer,
818 S., Kurten, T., Hu, W. W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Xu, L., Ng, N. L., Guo, H. Y.,
819 Weber, R. J., Wild, R. J., Brown, S. S., Koss, A., de Gouw, J., Olson, K., Goldstein, A. H., Seco, R., Kim,
820 S., McAvey, K., Shepson, P. B., Starn, T., Baumann, K., Edgerton, E. S., Liu, J. M., Shilling, J. E., Miller,
821 D. O., Brune, W., Schobesberger, S., D'Ambro, E. L., and Thornton, J. A.: Highly functionalized organic
822 nitrates in the southeast United States: Contribution to secondary organic aerosol and reactive nitrogen
823 budgets, *P Natl Acad Sci USA*, 113, 1516-1521, 10.1073/pnas.1508108113, 2016.

824 Liggio, J., and Li, S. M.: Reactive uptake of pinonaldehyde on acidic aerosols, *J Geophys Res-Atmos*, 111,
825 Artn D24303, 10.1029/2005jd006978, 2006.

826 Liggio, J., and Li, S. M.: Reversible and irreversible processing of biogenic olefins on acidic aerosols,
827 *Atmos Chem Phys*, 8, 2039-2055, DOI 10.5194/acp-8-2039-2008, 2008.

828 Liu, S., Ahlm, L., Day, D. A., Russell, L. M., Zhao, Y. L., Gentner, D. R., Weber, R. J., Goldstein, A. H.,
829 Jaoui, M., Offenberg, J. H., Kleindienst, T. E., Rubitschun, C., Surratt, J. D., Sheesley, R. J., and Scheller,
830 S.: Secondary organic aerosol formation from fossil fuel sources contribute majority of summertime organic
831 mass at Bakersfield, *J Geophys Res-Atmos*, 117, Artn D00v26, 10.1029/2012jd018170, 2012a.

832 Liu, S., Shilling, J. E., Song, C., Hiranuma, N., Zaveri, R. A., and Russell, L. M.: Hydrolysis of
833 Organonitrate Functional Groups in Aerosol Particles, *Aerosol Sci Tech*, 46, 1359-1369,
834 10.1080/02786826.2012.716175, 2012b.

835 Lopez-Hilfiker, F. D., Mohr, C., Ehn, M., Rubach, F., Kleist, E., Wildt, J., Mentel, T. F., Lutz, A., Hallquist,
836 M., Worsnop, D., and Thornton, J. A.: A novel method for online analysis of gas and particle composition:

837 description and evaluation of a Filter Inlet for Gases and AEROSols (FIGAERO), *Atmos Meas Tech*, 7,
838 983-1001, 10.5194/amt-7-983-2014, 2014.

839 Loza, C. L., Chan, A. W. H., Galloway, M. M., Keutsch, F. N., Flagan, R. C., and Seinfeld, J. H.:
840 Characterization of Vapor Wall Loss in Laboratory Chambers, *Environ Sci Technol*, 44, 5074-5078,
841 10.1021/es100727v, 2010.

842 Matsunaga, A., and Ziemann, P. J.: Gas-Wall Partitioning of Organic Compounds in a Teflon Film Chamber
843 and Potential Effects on Reaction Product and Aerosol Yield Measurements, *Aerosol Sci Tech*, 44, 881-
844 892, 10.1080/02786826.2010.501044, 2010.

845 Mcvay, R. C., Cappa, C. D., and Seinfeld, J. H.: Vapor-Wall Deposition in Chambers: Theoretical
846 Considerations, *Environ Sci Technol*, 48, 10251-10258, 10.1021/es502170j, 2014.

847 Muller, J. F., Peeters, J., and Stavrou, T.: Fast photolysis of carbonyl nitrates from isoprene, *Atmos Chem*
848 *Phys*, 14, 2497-2508, 10.5194/acp-14-2497-2014, 2014.

849 Nah, T., Sanchez, J., Boyd, C. M., and Ng, N. L.: Photochemical Aging of alpha-pinene and beta-pinene
850 Secondary Organic Aerosol formed from Nitrate Radical Oxidation, *Environ Sci Technol*, 50, 222-231,
851 10.1021/acs.est.5b04594, 2016.

852 Ng, N. L., Herndon, S. C., Trimborn, A., Canagaratna, M. R., Croteau, P. L., Onasch, T. B., Sueper, D.,
853 Worsnop, D. R., Zhang, Q., Sun, Y. L., and Jayne, J. T.: An Aerosol Chemical Speciation Monitor (ACSM)
854 for Routine Monitoring of the Composition and Mass Concentrations of Ambient Aerosol, *Aerosol Sci*
855 *Tech*, 45, 780-794, Pii 934555189, 10.1080/02786826.2011.560211, 2011.

856 Ng, N. L., Brown, S. S., Archibald, A. T., Atlas, E., Cohen, R. C., Crowley, J. N., Day, D. A., Donahue, N.
857 M., Fry, J. L., Fuchs, H., Griffin, R. J., Guzman, M. I., Herrmann, H., Hodzic, A., Iinuma, Y., Jimenez, J.
858 L., Kiendler-Scharr, A., Lee, B. H., Luecken, D. J., Mao, J. Q., McLaren, R., Mutzel, A., Osthoff, H. D.,
859 Ouyang, B., Picquet-Varrault, B., Platt, U., Pye, H. O. T., Rudich, Y., Schwantes, R. H., Shiraiwa, M.,
860 Stutz, J., Thornton, J. A., Tilgner, A., Williams, B. J., and Zaveri, R. A.: Nitrate radicals and biogenic
861 volatile organic compounds: oxidation, mechanisms, and organic aerosol, *Atmos Chem Phys*, 17, 2103-
862 2162, 10.5194/acp-17-2103-2017, 2017.

863 Nguyen, T. B., Crounse, J. D., Teng, A. P., Clair, J. M. S., Paulot, F., Wolfe, G. M., and Wennberg, P. O.:
864 Rapid deposition of oxidized biogenic compounds to a temperate forest, *P Natl Acad Sci USA*, 112, E392-
865 E401, 10.1073/pnas.1418702112, 2015.

866 Noziere, B., Barnes, I., and Becker, K. H.: Product study and mechanisms of the reactions of alpha-pinene
867 and of pinonaldehyde with OH radicals, *J Geophys Res-Atmos*, 104, 23645-23656, Doi
868 10.1029/1999jd900778, 1999.

869 Orlando, J. J., and Tyndall, G. S.: Laboratory studies of organic peroxy radical chemistry: an overview with
870 emphasis on recent issues of atmospheric significance, *Chem Soc Rev*, 41, 6294-6317,
871 10.1039/c2cs35166h, 2012.

872 Pankow, J. F.: An Absorption-Model of Gas-Particle Partitioning of Organic-Compounds in the
873 Atmosphere, *Atmos Environ*, 28, 185-188, Doi 10.1016/1352-2310(94)90093-0, 1994.

874 Pankow, J. F., and Asher, W. E.: SIMPOL.1: a simple group contribution method for predicting vapor
875 pressures and enthalpies of vaporization of multifunctional organic compounds, *Atmos Chem Phys*, 8,
876 2773-2796, DOI 10.5194/acp-8-2773-2008, 2008.

877 Perring, A. E., Pusede, S. E., and Cohen, R. C.: An Observational Perspective on the Atmospheric Impacts
878 of Alkyl and Multifunctional Nitrates on Ozone and Secondary Organic Aerosol, *Chem Rev*, 113, 5848-
879 5870, 10.1021/cr300520x, 2013.

880 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud
881 condensation nucleus activity, *Atmos Chem Phys*, 7, 1961-1971, DOI 10.5194/acp-7-1961-2007, 2007.

882 Pye, H. O. T., Luecken, D. J., Xu, L., Boyd, C. M., Ng, N. L., Baker, K. R., Ayres, B. R., Bash, J. O.,
883 Baumann, K., Carter, W. P. L., Edgerton, E., Fry, J. L., Hutzell, W. T., Schwede, D. B., and Shepson, P.
884 B.: Modeling the Current and Future Roles of Particulate Organic Nitrates in the Southeastern United States,
885 *Environ Sci Technol*, 49, 14195-14203, 10.1021/acs.est.5b03738, 2015.

886 Pye, H. O. T., D'Ambro, E. L., Lee, B., Schobesberger, S., Takeuchi, M., Zhao, Y., Lopez-Hilfiker, F., Liu,
887 J. M., Shilling, J. E., Xing, J., Mathur, R., Middlebrook, A. M., Liao, J., Welti, A., Graus, M., Warneke, C.,
888 de Gouw, J. A., Holloway, J. S., Ryerson, T. B., Pollack, I. B., and Thornton, J. A.: Anthropogenic
889 enhancements to production of highly oxygenated molecules from autoxidation, *P Natl Acad Sci USA*, 116,
890 6641-6646, 10.1073/pnas.1810774116, 2019.

891 Rindelaub, J. D., McAvey, K. M., and Shepson, P. B.: The photochemical production of organic nitrates
892 from alpha-pinene and loss via acid-dependent particle phase hydrolysis, *Atmos Environ*, 100, 193-201,
893 2015.

894 Rindelaub, J. D., Borca, C. H., Hostetler, M. A., Slade, J. H., Lipton, M. A., Slipchenko, L. V., and Shepson,
895 P. B.: The acid-catalyzed hydrolysis of an alpha-pinene-derived organic nitrate: kinetics, products, reaction
896 mechanisms, and atmospheric impact, *Atmos Chem Phys*, 16, 15425-15432, 10.5194/acp-16-15425-2016,
897 2016.

898 Rollins, A. W., Browne, E. C., Min, K. E., Pusede, S. E., Wooldridge, P. J., Gentner, D. R., Goldstein, A.
899 H., Liu, S., Day, D. A., Russell, L. M., and Cohen, R. C.: Evidence for NO_x Control over Nighttime SOA
900 Formation, *Science*, 337, 1210-1212, 10.1126/science.1221520, 2012.

901 Rollins, A. W., Pusede, S., Wooldridge, P., Min, K. E., Gentner, D. R., Goldstein, A. H., Liu, S., Day, D.
902 A., Russell, L. M., Rubitschun, C. L., Surratt, J. D., and Cohen, R. C.: Gas/particle partitioning of total
903 alkyl nitrates observed with TD-LIF in Bakersfield, *J Geophys Res-Atmos*, 118, 6651-6662,
904 10.1002/jgrd.50522, 2013.

905 Romonosky, D. E., Li, Y., Shiraiwa, M., Laskin, A., Laskin, J., and Nizkorodov, S. A.: Aqueous
906 Photochemistry of Secondary Organic Aerosol of alpha-Pinene and alpha-Humulene Oxidized with Ozone,
907 Hydroxyl Radical, and Nitrate Radical, *J Phys Chem A*, 121, 1298-1309, 10.1021/acs.jpca.6b10900, 2017.

908 Ruggeri, G., Bernhard, F. A., Henderson, B. H., and Takahama, S.: Model-measurement comparison of
909 functional group abundance in alpha-pinene and 1,3,5-trimethylbenzene secondary organic aerosol
910 formation, *Atmos Chem Phys*, 16, 8729-8747, 10.5194/acp-16-8729-2016, 2016.

911 Russell, L. M., Bahadur, R., and Ziemann, P. J.: Identifying organic aerosol sources by comparing
912 functional group composition in chamber and atmospheric particles, *P Natl Acad Sci USA*, 108, 3516-3521,
913 10.1073/pnas.1006461108, 2011.

914 Sanchez, J., Tanner, D. J., Chen, D. X., Huey, L. G., and Ng, N. L.: A new technique for the direct detection
915 of HO₂ radicals using bromide chemical ionization mass spectrometry (Br-CIMS): initial characterization,
916 *Atmos Meas Tech*, 9, 3851-3861, 10.5194/amt-9-3851-2016, 2016.

917 Saunders, S. M., Jenkin, M. E., Derwent, R. G., and Pilling, M. J.: Protocol for the development of the
918 Master Chemical Mechanism, MCM v3 (Part A): tropospheric degradation of non-aromatic volatile organic
919 compounds, *Atmos Chem Phys*, 3, 161-180, DOI 10.5194/acp-3-161-2003, 2003.

920 Schwantes, R. H., Teng, A. P., Nguyen, T. B., Coggon, M. M., Crouse, J. D., St Clair, J. M., Zhang, X.,
921 Schilling, K. A., Seinfeld, J. H., and Wennberg, P. O.: Isoprene NO₃ Oxidation Products from the RO₂ +
922 HO₂ Pathway, *J Phys Chem A*, 119, 10158-10171, 10.1021/acs.jpca.5b06355, 2015.

923 Seinfeld, J. H., and Pandis, S. N.: *Atmospheric chemistry and physics : from air pollution to climate change*,
924 2016.

925 Shilling, J. E., Zaveri, R. A., Fast, J. D., Kleinman, L., Alexander, M. L., Canagaratna, M. R., Fortner, E.,
926 Hubbe, J. M., Jayne, J. T., Sedlacek, A., Setyan, A., Springston, S., Worsnop, D. R., and Zhang, Q.:
927 Enhanced SOA formation from mixed anthropogenic and biogenic emissions during the CARES campaign,
928 *Atmos Chem Phys*, 13, 2091-2113, 10.5194/acp-13-2091-2013, 2013.

929 Shrivastava, M., Andreae, M. O., Artaxo, P., Barbosa, H. M. J., Berg, L. K., Brito, J., Ching, J., Easter, R.
930 C., Fan, J. W., Fast, J. D., Feng, Z., Fuentes, J. D., Glasius, M., Goldstein, A. H., Alves, E. G., Gomes, H.,
931 Gu, D., Guenther, A., Jathar, S. H., Kim, S., Liu, Y., Lou, S. J., Martin, S. T., McNeill, V. F., Medeiros,
932 A., de Sa, S. S., Shilling, J. E., Springston, S. R., Souza, R. A. F., Thornton, J. A., Isaacman-VanWertz, G.,
933 Yee, L. D., Ynoue, R., Zaveri, R. A., Zelenyuk, A., and Zhao, C.: Urban pollution greatly enhances
934 formation of natural aerosols over the Amazon rainforest, *Nat Commun*, 10, ARTN 1046, 10.1038/s41467-
935 019-08909-4, 2019.

936 Slade, J. H., de Perre, C., Lee, L., and Shepson, P. B.: Nitrate radical oxidation of gamma-terpinene:
937 hydroxy nitrate, total organic nitrate, and secondary organic aerosol yields, *Atmos Chem Phys*, 17, 8635-
938 8650, 10.5194/acp-17-8635-2017, 2017.

939 Song, M., Marcolli, C., Krieger, U. K., Zuend, A., and Peter, T.: Liquid-liquid phase separation in aerosol
940 particles: Dependence on O:C, organic functionalities, and compositional complexity, *Geophys Res Lett*,
941 39, 10.1029/2012gl052807, 2012.

942 Spittler, M., Barnes, I., Bejan, I., Brockmann, K. J., Benter, T., and Wirtz, K.: Reactions of NO₃ radicals
943 with limonene and alpha-pinene: Product and SOA formation, *Atmos Environ*, 40, S116-S127,
944 10.1016/j.atmosenv.2005.09.093, 2006.

945 Spracklen, D. V., Jimenez, J. L., Carslaw, K. S., Worsnop, D. R., Evans, M. J., Mann, G. W., Zhang, Q.,
946 Canagaratna, M. R., Allan, J., Coe, H., McFiggans, G., Rap, A., and Forster, P.: Aerosol mass spectrometer
947 constraint on the global secondary organic aerosol budget, *Atmos Chem Phys*, 11, 12109-12136,
948 10.5194/acp-11-12109-2011, 2011.

949 Stark, H., Yatavelli, R. L. N., Thompson, S. L., Kang, H., Krechmer, J. E., Kimmel, J. R., Palm, B. B., Hu,
950 W. W., Hayes, P. L., Day, D. A., Campuzano-Jost, P., Canagaratna, M. R., Jayne, J. T., Worsnop, D. R.,
951 and Jimenez, J. L.: Impact of Thermal Decomposition on Thermal Desorption Instruments: Advantage of
952 Thermogram Analysis for Quantifying Volatility Distributions of Organic Species, *Environ Sci Technol*,
953 51, 8491-8500, 10.1021/acs.est.7b00160, 2017.

954 Surratt, J. D., Gomez-Gonzalez, Y., Chan, A. W. H., Vermeylen, R., Shahgholi, M., Kleindienst, T. E.,
955 Edney, E. O., Offenberg, J. H., Lewandowski, M., Jaoui, M., Maenhaut, W., Claeys, M., Flagan, R. C., and
956 Seinfeld, J. H.: Organosulfate formation in biogenic secondary organic aerosol, *J Phys Chem A*, 112, 8345-
957 8378, 10.1021/jp802310p, 2008.

958 Takeuchi, M., and Ng, N. L.: Organic Nitrates and Secondary Organic Aerosol (SOA) Formation from
959 Oxidation of Biogenic Volatile Organic Compounds, *Acs Sym Ser*, 1299, 105-125, 2018.

960 Wangberg, I., Barnes, I., and Becker, K. H.: Product and mechanistic study of the reaction of NO₃ radicals
961 with alpha-pinene, *Environ Sci Technol*, 31, 2130-2135, DOI 10.1021/es960958n, 1997.

962 Wayne, R. P., Barnes, I., Biggs, P., Burrows, J. P., Canosamas, C. E., Hjorth, J., Lebras, G., Moortgat, G.
963 K., Perner, D., Poulet, G., Restelli, G., and Sidebottom, H.: The Nitrate Radical - Physics, Chemistry, and
964 the Atmosphere, *Atmos Environ a-Gen*, 25, 1-203, Doi 10.1016/0960-1686(91)90192-A, 1991.

965 Weber, R. J., Sullivan, A. P., Peltier, R. E., Russell, A., Yan, B., Zheng, M., de Gouw, J., Warneke, C.,
966 Brock, C., Holloway, J. S., Atlas, E. L., and Edgerton, E.: A study of secondary organic aerosol formation
967 in the anthropogenic-influenced southeastern United States, *J Geophys Res-Atmos*, 112, Artn D13302,
968 10.1029/2007jd008408, 2007.

969 Xu, L., Guo, H. Y., Boyd, C. M., Klein, M., Bougiatioti, A., Cerully, K. M., Hite, J. R., Isaacman-VanWertz,
970 G., Kreisberg, N. M., Knote, C., Olson, K., Koss, A., Goldstein, A. H., Hering, S. V., de Gouw, J., Baumann,
971 K., Lee, S. H., Nenes, A., Weber, R. J., and Ng, N. L.: Effects of anthropogenic emissions on aerosol
972 formation from isoprene and monoterpenes in the southeastern United States (vol 112, pg 37, 2015), *P Natl*
973 *Acad Sci USA*, 112, E4509-E4509, 10.1073/pnas.1512279112, 2015a.

974 Xu, L., Suresh, S., Guo, H., Weber, R. J., and Ng, N. L.: Aerosol characterization over the southeastern
975 United States using high-resolution aerosol mass spectrometry: spatial and seasonal variation of aerosol
976 composition and sources with a focus on organic nitrates, *Atmos Chem Phys*, 15, 7307-7336, 10.5194/acp-
977 15-7307-2015, 2015b.

978 Xu, L., Pye, H. O. T., He, J., Chen, Y. L., Murphy, B. N., and Ng, N. L.: Experimental and model estimates
979 of the contributions from biogenic monoterpenes and sesquiterpenes to secondary organic aerosol in the
980 southeastern United States, *Atmos Chem Phys*, 18, 12613-12637, 10.5194/acp-18-12613-2018, 2018a.

981 Xu, L., Moller, K. H., Crouse, J. D., Otkjwr, R. V., Kjaergaard, H. G., and Wennberg, P. O.: Unimolecular
982 Reactions of Peroxy Radicals Formed in the Oxidation of alpha-Pinene and beta-Pinene by Hydroxyl
983 Radicals, *J Phys Chem A*, 123, 1661-1674, 10.1021/acs.jpca.8b11726, 2019.

984 Xu, W., Lambe, A., Silva, P., Hu, W. W., Onasch, T., Williams, L., Croteau, P., Zhang, X., Renbaum-
985 Wolff, L., Fortner, E., Jimenez, J. L., Jayne, J., Worsnop, D., and Canagaratna, M.: Laboratory evaluation
986 of species-dependent relative ionization efficiencies in the Aerodyne Aerosol Mass Spectrometer, *Aerosol*
987 *Sci Tech*, 52, 626-641, 10.1080/02786826.2018.1439570, 2018b.

988 Zare, A., Romer, P. S., Nguyen, T., Keutsch, F. N., Skog, K., and Cohen, R. C.: A comprehensive organic
989 nitrate chemistry: insights into the lifetime of atmospheric organic nitrates, *Atmos Chem Phys*, 18, 15419-
990 15436, 10.5194/acp-18-15419-2018, 2018.

991 Zhang, H. F., Yee, L. D., Lee, B. H., Curtis, M. P., Worton, D. R., Isaacman-VanWertz, G., Offenberg, J.
992 H., Lewandowski, M., Kleindienst, T. E., Beaver, M. R., Holder, A. L., Lonneman, W. A., Docherty, K. S.,
993 Jaoui, M., Pye, H. O. T., Hu, W. W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Guo, H. Y., Weber,
994 R. J., de Gouw, J., Koss, A. R., Edgerton, E. S., Brune, W., Mohr, C., Lopez-Hilfiker, F. D., Lutz, A.,
995 Kreisberg, N. M., Spielman, S. R., Hering, S. V., Wilson, K. R., Thornton, J. A., and Goldstein, A. H.:
996 Monoterpenes are the largest source of summertime organic aerosol in the southeastern United States, *P*
997 *Natl Acad Sci USA*, 115, 2038-2043, 10.1073/pnas.1717513115, 2018.

- 998 Zhang, X., Cappa, C. D., Jathar, S. H., Mcvay, R. C., Ensberg, J. J., Kleeman, M. J., and Seinfeld, J. H.:
999 Influence of vapor wall loss in laboratory chambers on yields of secondary organic aerosol, P Natl Acad
1000 Sci USA, 111, 5802-5807, 10.1073/pnas.1404727111, 2014.
- 1001 Zhang, X., Schwantes, R. H., McVay, R. C., Lignell, H., Coggon, M. M., Flagan, R. C., and Seinfeld, J. H.:
1002 Vapor wall deposition in Teflon chambers, Atmos Chem Phys, 15, 4197-4214, 10.5194/acp-15-4197-2015,
1003 2015.
- 1004 Ziemann, P. J., and Atkinson, R.: Kinetics, products, and mechanisms of secondary organic aerosol
1005 formation, Chem Soc Rev, 41, 6582-6605, 10.1039/c2cs35122f, 2012.
- 1006

1 **Supporting Information for**

2

3 **Chemical Composition and Hydrolysis of Organic Nitrate Aerosol formed from Hydroxyl and**
4 **Nitrate Radical Oxidation of α -pinene and β -pinene**

5

6 Masayuki Takeuchi¹ and Nga L. Ng^{2,3*}

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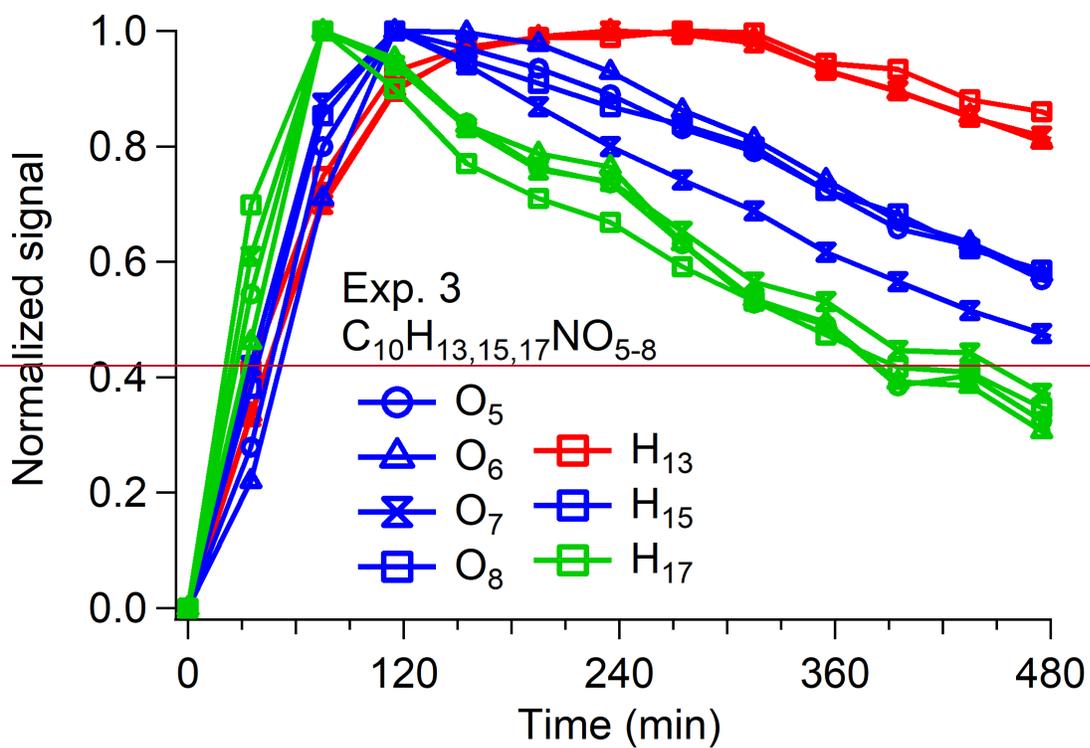
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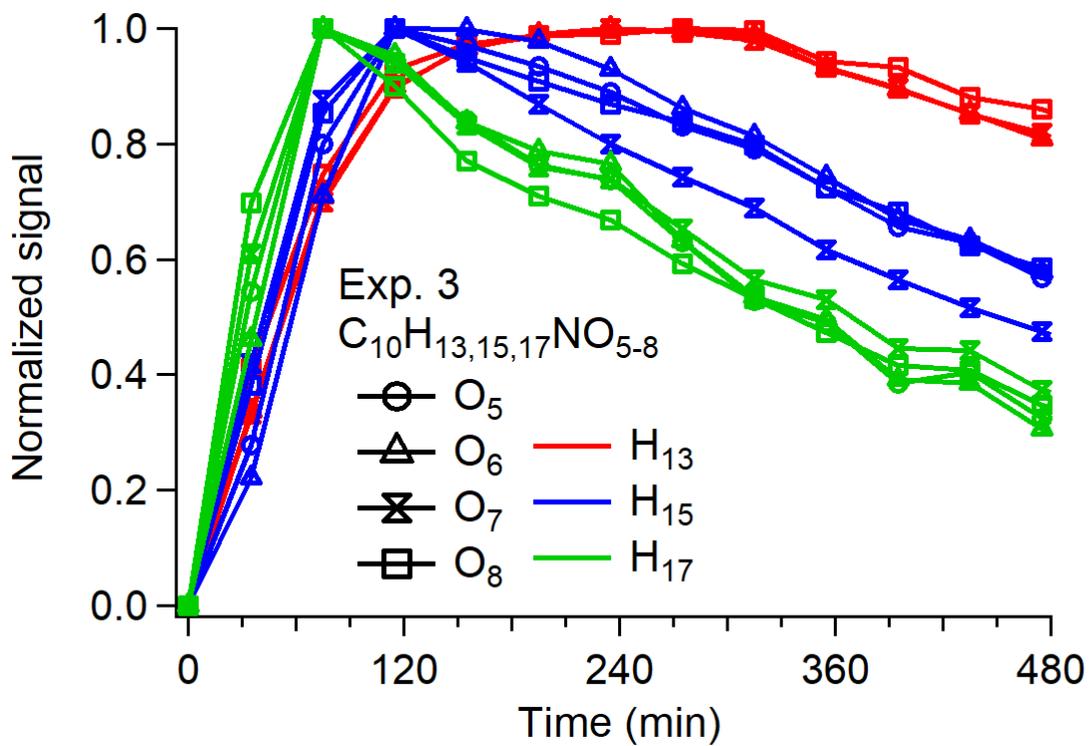
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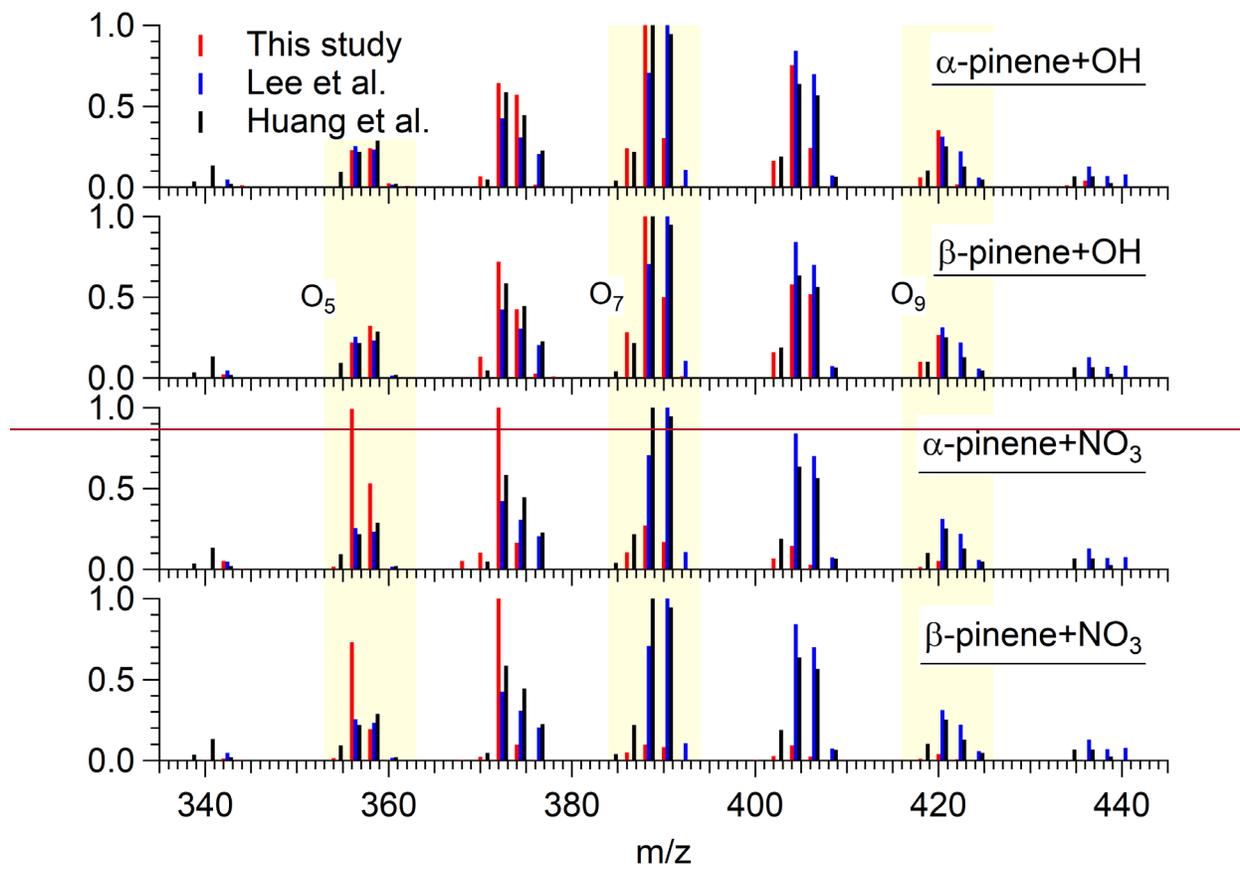
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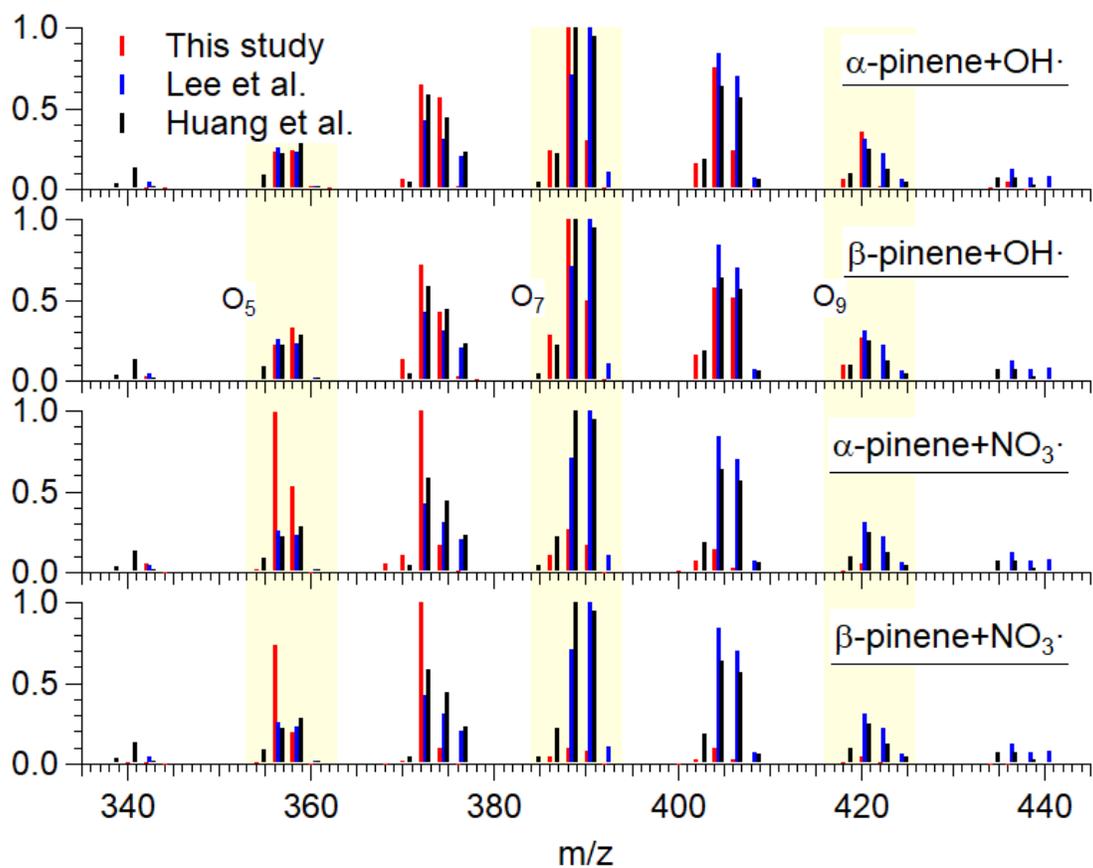


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18 Figure S1. FIGAERO-HR-ToF-I-CIMS time-series data of select organic nitrate aerosols ~~(i.e.,~~
19 ~~$C_{10}H_{13,15,17}NO_{5-8}$)~~. All combinations of $C_{10}H_{13,15,17}NO_{5-8}$ with different oxygen and nitrogen numbers are
20 shown here. Different symbols correspond to compounds with different oxygen numbers, while different
21 colors correspond to compounds with different nitrogen numbers. Shown are the data from Exp. 3 (α -
22 pinene+OH₂), though very similar trends are observed for β -pinene+OH₂ from Exp. 6.

23

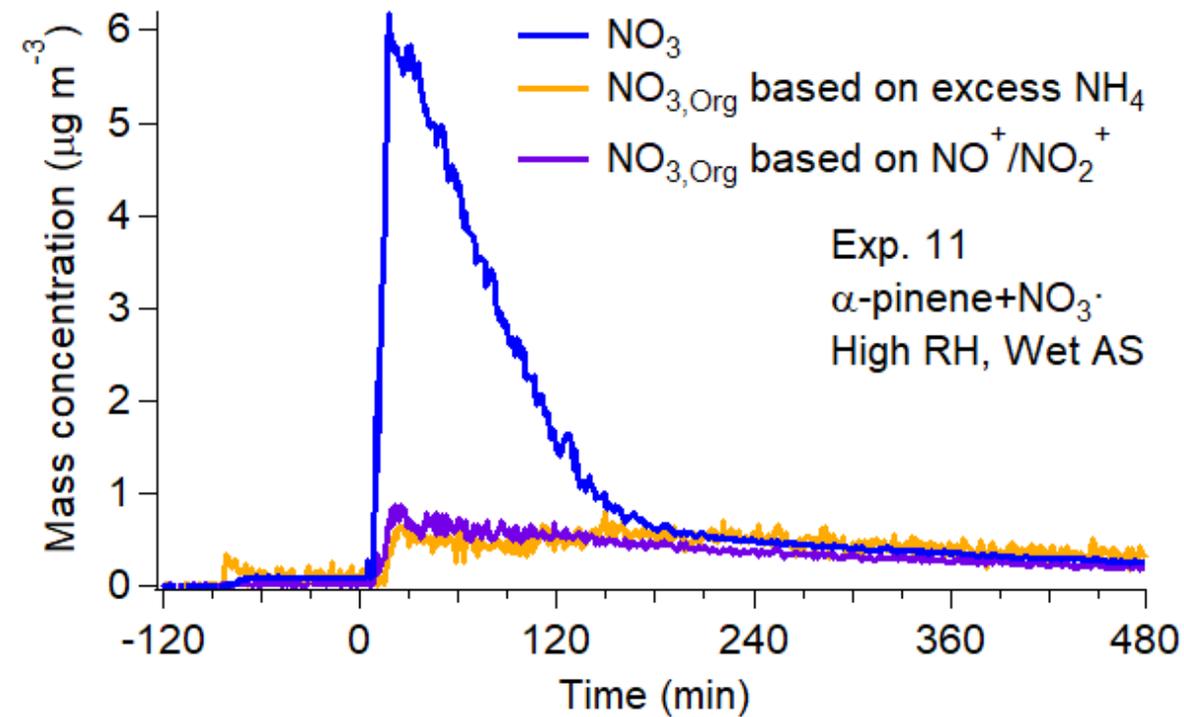
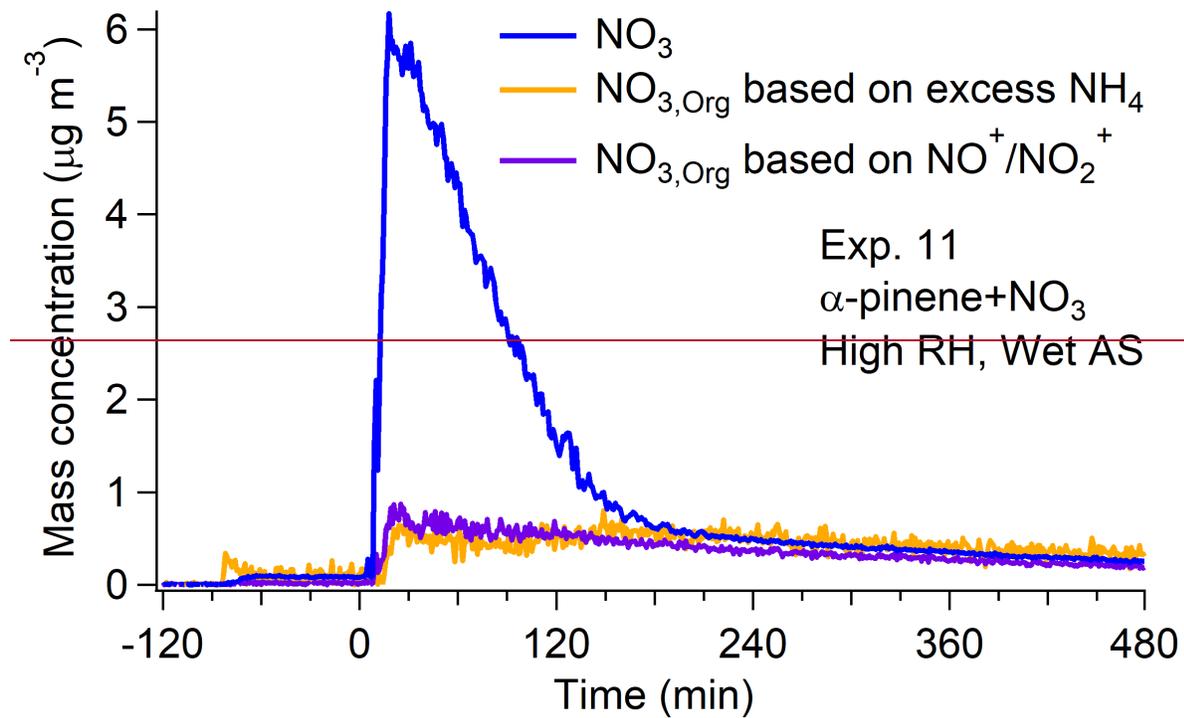




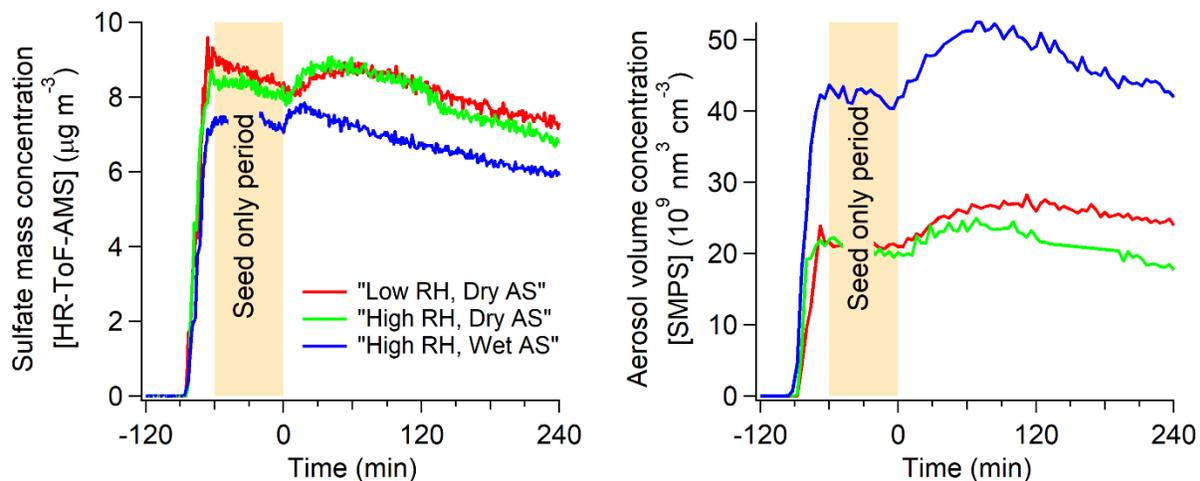
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26 Figure S2. Comparison of FIGAERO-HR-ToF-I-CIMS mass spectra of organic nitrate aerosols (i.e.,
 27 $C_{10}H_{11,13,15,17,19,21}NO_{4-11}$) with ambient measurement data (Lee et al., 2016; Huang et al., 2019). All mass
 28 spectra are normalized to the maximum signal.

29



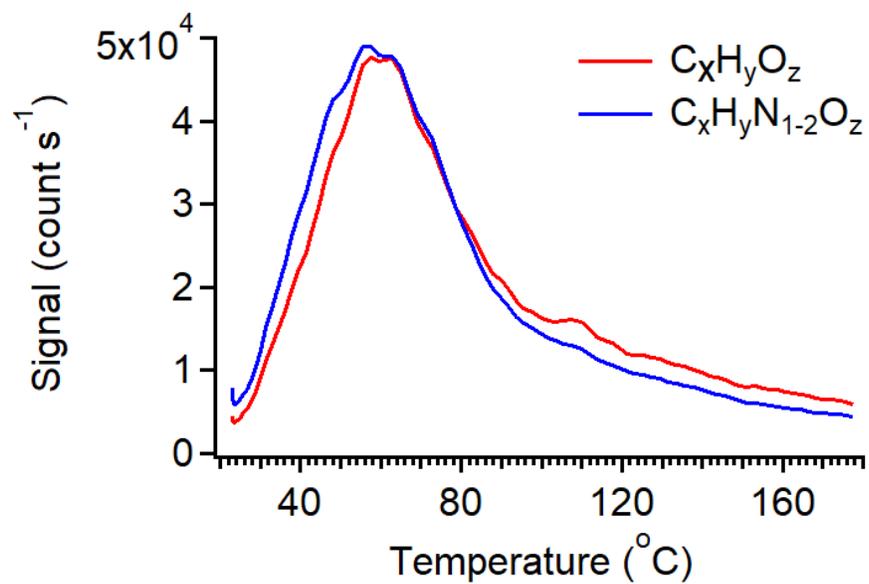
- 33 Figure S3. Concentrations of NO_3 and $\text{NO}_{3,\text{Org}}$ derived from two independent methods. “Excess NH_4 ”
- 34 refers to a method to apportion the inorganic nitrate contribution to NO_3 based on the increase in NH_4 and
- 35 “ $\text{NO}^+/\text{NO}_2^+$ ” refers to the method based on its ratio (Farmer et al., 2010).



36

37 Figure S4. HR-ToF-AMS time-series data of SO_4 and SMPS aerosol volume concentration in Exp. 3 (low
 38 RH, dry AS), Exp. 4 (high RH, dry AS), and Exp. 5 (high RH, wet AS). A significant difference in the
 39 volume concentration among the experiments indicates the deliquesced nature in Exp. 5 (high RH, wet AS)
 40 and the effloresced nature of seed particles in Exp. 4 (high RH, dry AS).

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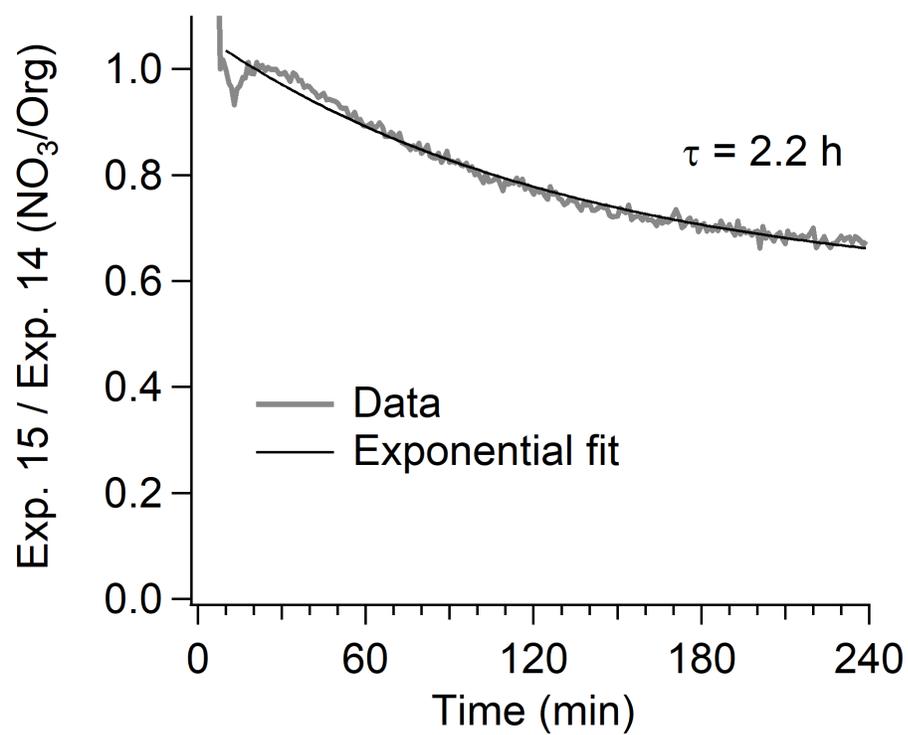
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43 Figure S5. Thermal desorption profiles of C_xH_yO_z and C_xH_yN₁₋₂O_z in FIGAERO-HR-ToF-I-CIMS from

44 Exp. 3. Data points are averages of three desorption cycles around the peak SOA growth.

45

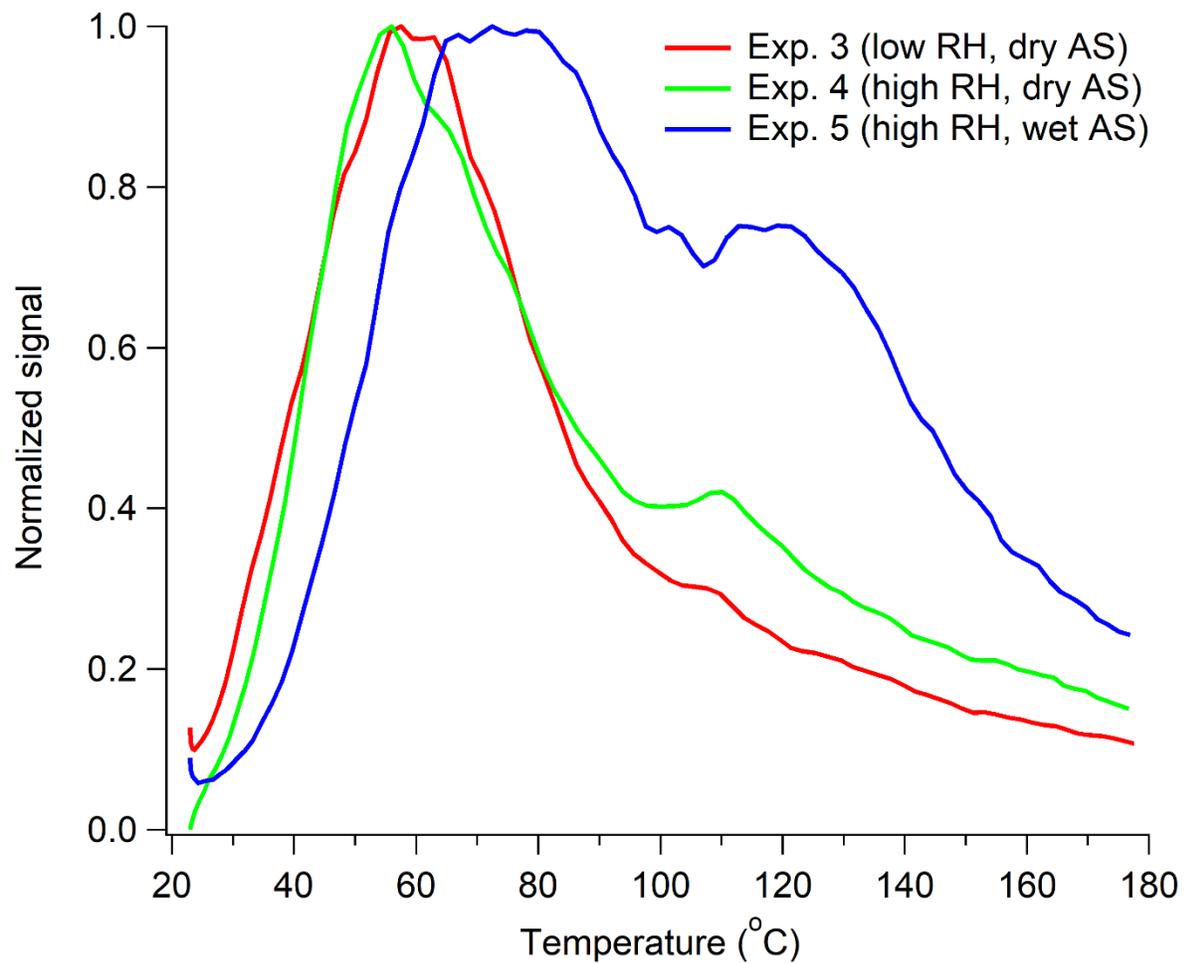
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48 Figure S56. Time-series data of NO₃/Org in Exp. 15 normalized to that in Exp. 14 based on the method
49 present in prior study (Boyd et al., 2015).

50



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52 Figure S67. Average thermogram of organic aerosol (i.e., organic nitrate and non-nitrated organic
53 aerosol) at the peak of SOA growth in Exp. 3-5. Signals are normalized to the maximum signal for each
54 experiment.

55 Table S1. Summary of R_{AN} and R_{ON} applied to derive the concentration of $NO_{3,Org}$. R_{AN} from 300 nm
56 ammonium nitrate particle at the nearest date is used unless noted otherwise. R_{ON}/R_{AN} used for high RH
57 experiments are taken from the same experiment but under low RH experiments.

ID	R_{AN}	R_{ON}	R_{ON}/R_{AN}
1	2.87	6.25	2.18
2	2.87	6.25	2.18
3	2.80	5.75	2.05
4	2.80	5.75	2.05
5	2.88	5.91	2.05
6	3.06	5.02	1.64
7	3.06	5.02	1.64
8	3.24	8.64	2.67
9	2.87	7.66	2.67
10	2.70	8.24	3.05
11	2.80	8.55	3.05
12	2.80	8.60	3.07
13	2.88	8.85	3.07
14	2.87	7.13	2.48
15	2.56 ^a	6.36	2.48

58 ^a R_{AN} is taken from 30-min average of NO^+/NO_2^+ during seed only period as the R_{AN} from the calibration at
59 the nearest date is slightly higher.

60

61

62 Table S2. Fraction of pON based on the number of carbon in each system.

System	C _{≤9} pON	C ₁₀ pON	C ₁₁₋₁₉ pON	C ₂₀ pON
α-pinene+OH ₂	47 %	52 %	1 %	0 %
β-pinene+OH ₂	58 %	41 %	1 %	0 %
α-pinene+NO ₃	11 %	26 %	8 %	54 %
β-pinene+NO ₃	10 %	58 %	4 %	28 %

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References

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66 Boyd, C. M., Sanchez, J., Xu, L., Eugene, A. J., Nah, T., Tuet, W. Y., Guzman, M. I., and Ng, N. L.:
67 Secondary organic aerosol formation from the beta-pinene+NO₃ system: effect of humidity and peroxy
68 radical fate, *Atmos Chem Phys*, 15, 7497-7522, 10.5194/acp-15-7497-2015, 2015.
- 69 Farmer, D. K., Matsunaga, A., Docherty, K. S., Surratt, J. D., Seinfeld, J. H., Ziemann, P. J., and Jimenez,
70 J. L.: Response of an aerosol mass spectrometer to organonitrates and organosulfates and implications for
71 atmospheric chemistry, *P Natl Acad Sci USA*, 107, 6670-6675, 10.1073/pnas.0912340107, 2010.
- 72 Huang, W., Saathoff, H., Shen, X. L., Ramisetty, R., Leisner, T., and Mohr, C.: Chemical Characterization
73 of Highly Functionalized Organonitrates Contributing to Night-Time Organic Aerosol Mass Loadings and
74 Particle Growth, *Environ Sci Technol*, 53, 1165-1174, 10.1021/acs.est.8b05826, 2019.
- 75 Lee, B. H., Mohr, C., Lopez-Hilfiker, F. D., Lutz, A., Hallquist, M., Lee, L., Romer, P., Cohen, R. C., Iyer,
76 S., Kurten, T., Hu, W. W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Xu, L., Ng, N. L., Guo, H. Y.,
77 Weber, R. J., Wild, R. J., Brown, S. S., Koss, A., de Gouw, J., Olson, K., Goldstein, A. H., Seco, R., Kim,
78 S., McAvey, K., Shepson, P. B., Starn, T., Baumann, K., Edgerton, E. S., Liu, J. M., Shilling, J. E., Miller,
79 D. O., Brune, W., Schobesberger, S., D'Ambro, E. L., and Thornton, J. A.: Highly functionalized organic
80 nitrates in the southeast United States: Contribution to secondary organic aerosol and reactive nitrogen
81 budgets, *P Natl Acad Sci USA*, 113, 1516-1521, 10.1073/pnas.1508108113, 2016.

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