

Response to the Comments from Referee #1

General comments: Ozone (O_3), as a criteria air pollutant, is attracting increasing concerns in China due to the rapid rise in concentrations across the country. This paper studied the O_3 pollution in an economically boomed region of China (PRD in southern China), in terms of the meteorological impacts, local contribution and regional transport. The statistical methods were used, with interesting and meaningful results being reported. Basically, the decadal changes of O_3 in PRD were well explained, except for some aspects where further clarifications or reorganizations are needed. Though some of the findings were already known knowledge, the paper further consolidate our understandings and fully demonstrated its value in future O_3 pollution control in this region. Thus, the paper is recommended to be accepted after the following problems are addressed.

1. First, I do not quite agree with the authors' statement of the "conceptual model". Generally, a conceptual model is established based on some phenomena, and is further verified by the results. In this study, I would like to suggest the authors to replace the "conceptual model" with the "discussions" on the results, because I did not see the verification of the "model".

Response: According to Pun and Seigneur (1999), a conceptual model is a qualitative compilation of the physical and chemical processes that govern the formation of pollutants, which, to the extent possible, is supported by quantitative information. Conceptual model helps to elucidate pollutant contributions from local and non-local sources and their temporal and spatial variations, and the underlying reasons governing such variations. In this study, we use information about ozone precursor (VOCs and NO_x) emissions and their changing trends, ozone formation regimes, and the monsoonal and micro-scale synoptic conditions over different sub-regions of the Pearl River Delta to explain the spatiotemporal variations of modeling object (locally formed O_3 with meteorological adjustment). Therefore, we believe our analysis approach could be called a 'conceptual model' and would like to keep using it. Good interpretation on the modeling object (locally formed O_3 with meteorological adjustment) both under general conditions and during episodes serves as verification of the applicability of the conceptual model.

2. Second, the discussions on the O_3 episodes were relatively weak and the trend analyses did not seem to be appropriate for O_3 episodes, due to the limited and inconsistent number of episode days in every years. This section needs to be reorganized and some discussions should be clarified or corrected.

Response: I don't know why the reviewer thinks the trend analysis did not seem to be appropriate for O_3 episodes. Yes the number of episodic days are different across the 11 years, such a difference can be largely accounted and removed by meteorological adjustment. In this section, we add explanation on the slight changes of non-local emission contribution on ozone after 2014.

3. Third, the discussions on O_3 pollution in 2016 and 2017 look a bit weird, which should be reorganized.

Response: The logic here is that in previous sections we have examined the long-term trends of meteorological impacts, and O_3 contributions from non-local emissions and different factors of local emissions. Our statistical analysis framework also has the capability of quantitatively separating the impact of meteorology and local and non-local emissions. Actually we can do this

year by year. Considering the length of the paper, we decide to show this capability in the year of 2016 and 2017, the most recent two years in the study period.

4. Lastly, it will be good if the O₃ distribution, trend and the influencing factors can be discussed separately by seasons. In fact, the increases of springtime O₃ in PRD in recent years were striking, in contrast to the overall unchanged O₃ in summer and autumn. The paper would be more informative with the discussions on O₃ pollution in different seasons and a special focus on the season when O₃ is highest or increasing with the highest rates.

Response: The main point of this paper is to quantitatively examine the impacts of meteorology and precursor emissions from within and outside the PRD on the evolution of ozone during the past decade by using a statistical analysis framework combining meteorological adjustment and source apportionment and a conceptual model. Seasonal variation is indeed a very interesting issue but cannot be addressed in detail here due to the length of the paper. We will investigate it in our further study.

Specific comments:

1. Page 6, line 233-234. What were the reasons of the minor changes in O₃ due to emissions during 2011 - 2015, in contrast to the significant increase before 2011? Throughout the paper, the changes in the meteorological and artificial impacts, especially the turning points of O₃ variations, should be discussed.

Response: According to the PRD emission inventory developed by our research group (Figure S1, also provided as Figure 1 below, manuscript under preparation), increase in VOCs emission started to mitigate in 2011, while NO_x emissions showed significant reduction starting from 2013. As PRD is generally in a VOC-limited ozone formation regime, reduction in the magnitude of VOCs emission increase is likely responsible for the minor changes in O₃ during 2011-2015. This is added in lines 230-234 in the revised manuscript.

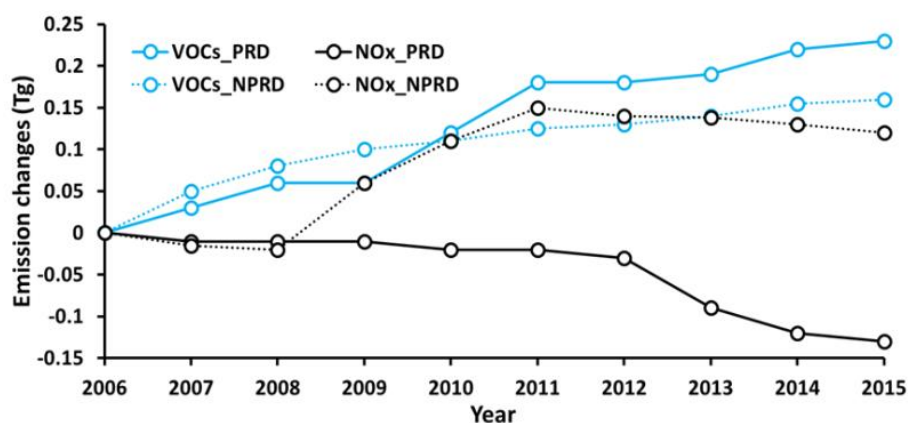


Fig 1 Emissions of NO_x and VOCs in the Pearl River Delta during 2006-2016

2. Page 6, lines 236-238. The strong statement must be evidenced. This statement is actually contradictory to the later finding that “meteorological adjustment does not alter ozone concentration much” (page 9, lines 362 – 363) on episode days. Please clarify.

Response: Agree with the reviewer that we cannot draw the conclusion that meteorological condition is “the most” important driving factor. We change it to “one of the most” in lines 243-

244 in the revised manuscript, as it can be clearly evidenced by Fig. 3(b) that ozone spikes were shortened after meteorological adjustment. This indicates that meteorology plays a leading role in generating ozone episodes. It should be noted that meteorological adjustment here is against the average meteorological condition during the entire period. In comparison, the meteorological adjustment in the episodic analysis is against the average meteorological condition during episode days. Although the average meteorological condition varies greatly in different years, that during episodes does not differ much in different years. Therefore, we conclude that “meteorological adjustment does not alter ozone concentration much” on episode days.

3. Page 6, lines 259 – 260. Please briefly explain why central and western PRD were the regions that were most sensitive to meteorological conditions in O₃ pollution?

Response: Generally, area with higher pollutant emission is more sensitive to changes in meteorological condition (Seo et al., 2014). Various studies have shown that central and western PRD is the area with the most intense O₃ precursor (VOCs and NO_x) emissions over the PRD (e.g. Zheng et al., 2009a) therefore is more sensitive to meteorological condition. This is added in lines 266-269 in the revised manuscript.

4. Page 7, line 267. “...and that most local emissions are concentrated in the central PRD area”. Add references to support the statement.

Response: We revise this statement as “...and that most local emissions are concentrated in the central and western PRD area”. Reference Zheng et al. (2009a) is added in the revision.

5. Page 7, lines 303 – 305. The reasons for picking the three sites should be given. It will be better to plot all the sites in the supplement.

Response: There is a north-south gradient in the spatial distributions of all three principal component loadings, as shown in Fig. 5a. Therefore, we select Tianhu (TH) in the north, Luhu (LH) in central, and Donghu (DH) in the south of PRD to study the long-term trend of ozone contributed by local and non-local emission sources in different areas. Ozone contributions at the other stations is added in Fig. S2 of the supplement.

6. Page 8, lines 349 – 351. Change “strengthening” to “constraining” or “restraining”. Why only VOCs should be controlled? Also, cutting VOCs emissions will not prevent the decrease of NO titration to O₃. The statements should be more accurate throughout.

Response: Here we are talking about “strengthening local VOC emission control”. There is a “control” at the end of the sentence, therefore we cannot change “strengthening” to “constraining”. In addition, NO titration of O₃ is dependent on VOC/NO_x ratio. If the ratio is lower, NO titration of O₃ near emission sources would be higher. Therefore, controlling VOCs emission to reduce VOC/NO_x ratio would to some extent enhance NO titration so as to reduce O₃ level.

7. Page 9, lines 377 – 378. What were the causes of levelling off and decrease of non-local contributions? As commented above, the changes are worth to be discussed, which may relate to the nationwide emission controls.

Response: At this time we are not aware what the exact reason for levelling off and decrease of non-local contribution after 2014. We look at VOCs and NO_x emission inventory over non-PRD area in Guangdong Province developed by our research group (Figure 1). NO_x emissions decreases and

VOCs emissions increases after 2014. As most of the non-PRD area in Guangdong may in a transitional ozone formation regime in an annual average, such a slight change in NO_x and VOC emissions lead to little changes in O₃ formation and transport into the PRD. We must emphasize that the above argument is simply our speculation and further modeling study is needed to explain such a phenomenon.

8. Page 11, lines 448 – 450. It is most likely that O₃ formation in the northeastern PRD became more limited by NO_x, however evidences should be provided to prove the shift of O₃ formation regime from VOC-limited to NO_x-limited in the southwestern.

Response: There is no publication reporting the shift in O₃ formation regime in southwestern PRD, as such a shift mainly occurs in recent years. However, we may speculate such a shift during pollution episodes considering the intense biogenic VOC emissions over southwestern rural area when temperature is high and solar radiation is strong. Jin and Holloway (2015) discovered that O₃ photochemistry is NO_x limited from April to October and transitional or mixed in other months using satellite observed HCHO/NO₂ ratio. Although there is large uncertainty by using HCHO/NO₂ to infer ozone formation regime, such a trend of shifting to NO_x-limited regime during ozone episodes is evidenced.

9. Page 11, lines 455 – 464. The discussions on O₃ episodes need to be deepened. For example, the winds were not always from the east during O₃ episodes, which in fact were from the northeast with low speeds in most cases under continental anticyclones, and from the northwest with the approaching of tropical cyclones. I do not think that the winds during O₃ episodes can be simplified as easterly, so did the other characteristics which were discussed as an integration in this paper.

Response: In comparison with general conditions, prevailing winds during episodes are more easterly (Figure S6, also provided as Figure 2 below, manuscript under preparation),. We are not saying that the winds were always from the east during ozone episodes. Slight shift to northeasterly or southeasterly would not change the conclusion from the conceptual model that southwestern PRD around DH, ZML and TJ stations is a sink region of ozone produced around the Pearl River Estuary, thereby having high ozone levels during episodes. Northwest wind by the approaching of tropical cyclones could bring ozone episodes to HK and Shenzhen as they are located downwind of central PRD. However, from an entire region point of view, northwest wind during ozone episodes is rather limited.

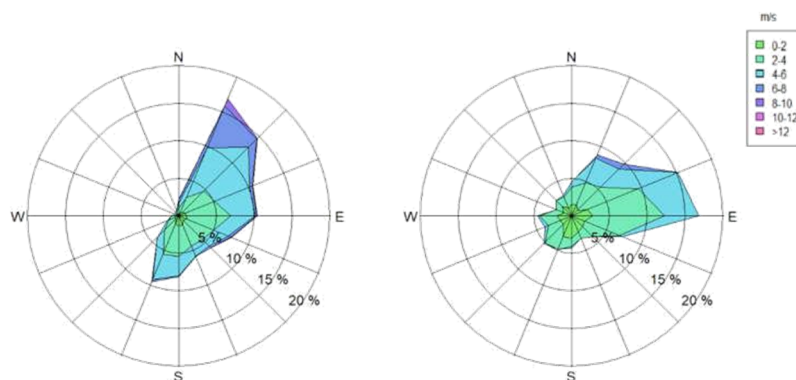


Fig 2. Wind rose under general conditions (left) and during ozone episodes (right) during 2007 and 2017 in the Pearl River Delta.

10. Page 13, lines 529 – 530. The term “optimal effective NO_x/VOC ratio” needs to be annotated. The optimal ratio in fact means highest O₃ production rates, which is the worst from the angle of O₃ pollution control.

Response: Here we are exactly talking about NO_x/VOC ratio with the highest O₃ concentration. This is added in lines 496-497 of the revised manuscript. Our previous publication (Ou et al., 2016) concluded that in the VOC-limited regime, O₃ reduction can be possible in short-term but cannot reduce O₃ into attainment. The only way is to control NO_x emission so as to change ozone formation regime from VOC-limited to NO_x-limited. In this process, O₃ concentration would increase slightly in the beginning, and drop significantly after bypassing the turning point with the optimal effective NO_x/VOC ratio that corresponds with the highest O₃ level.

References

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Response to the Comments from Referee #2

The manuscript by Yang et al. describes the quantitatively statistical analysis of meteorology and precursor emissions impacts on ozone evolution in PRD region. The strategy of the analysis is clear and the manuscript has been relatively well organized for such a topic. This study presents an innovative approach to quantify the contributions of meteorological factors as well as non-local precursor emissions for ozone concentration in PRD. Besides, this study reveals the special distribution differences of ozone increase contribution under general conditions and ozone episodes, and uses a model to explain the difference result from emission control. Furthermore, specific suggestions are given in order to acquire high ozone reduction efficiency. Therefore it merits to be published in ACP. However, more detailed explanations are expected to make this manuscript complete and more convincing. In the following, I had a number of specific comments for the authors' reference to address before publication.

Specific comments:

1. It is confusing for the final sentence at the second paragraph of section 3.1. How to conclude the meteorological condition is the most important driving factor for ozone episodes through Fig. 3b?

Response: Agree with the reviewer that we cannot draw the conclusion that meteorological condition is "the most" important driving factor. We change it to "one of the most" in lines 243-244 in the revised manuscript, as it can be clearly evidenced by Fig. 3b that ozone spikes were shortened after meteorological adjustment.

2. In Fig. 4a, what is the legend of the color contour?

Response: The legend of color contours in Fig. 4a refer to the isopleths of averaged ozone concentration ($\mu\text{g m}^{-3}$) both before and after meteorological adjustment, as illustrated in the caption of Fig. 4a.

3. In the first paragraph of section 3.3, considering the monsoon climate, why local emissions have distinct impacts while non-local emissions are consistent?

Response: Agree that the term 'consistent' here leads to misunderstanding. In fact, non-local emissions do pose different impacts to different areas of the PRD due much to deposition and local topography. However, the impacts at all areas are positive, which is different from the impact by local emissions which is positive in some directions while negative in others. To make it more accurate, we revise the sentence as "Considering PRD's monsoonal synoptic condition and that most local emissions are concentrated in the central PRD area, local ('within PRD') emissions tend to pose contrasting impacts to different sub-regions in different seasons while the impacts from non-local ('outside PRD') emissions in a relatively larger scale tend to be similar over the PRD.", as shown in lines 277-279 of the revised manuscript.

4. In section 3.3, I would like to suggest the authors to further explain how to distinguish the non-local emissions from the impact of meteorological condition, as the PC1 is associated with northeasterly wind and meteorology has been removed in meteorological adjustment part.

Response: As explained in section 2.2, meteorological adjustment is conducted against the base condition which is the average meteorological condition of the same calendar date throughout 11 years. Therefore, meteorological adjustment removes the inter-annual variation of meteorological condition, while still preserves seasonal variation and make it consistent in all years. Therefore, the monsoonal characteristics

(northerly or northeasterly winds in winter and southerly and southwesterly winds in summer) still remain and act as an important justification for PC determination.

5. What does the score mean in Fig. 5b and Fig. 6? I would like to suggest the authors add more explanations in figure caption for Fig. 5b and Fig. 6.

Response: In the EOF analysis, the score of a PC refers to the temporal variation of the loading of that PC. For example, in Fig. 5b, the score of PC1 shows the increasing contribution of the loading of PC1, which is determined as non-local emission contribution due to relatively similar loadings throughout the PRD.

6. As SSR plays an important role in meteorology factors, I wonder if it is necessary to add the correlation of SSR with PCs?

Response: The three PCs here refer to the impacts of local and non-local emissions on O₃. Therefore, we consider there is no relation between the impacts of emissions and meteorological factors, such as SSR.

7. In sections 3.3 and 3.5, long-term trends of local and non-local emission contributions on ozone is an interesting finding. The authors are suggested to provide indepth discussion the influencing factors shaping the trends rather than simply describing them, both under general conditions and during ozone episodes.

Response: We believe the long-term increasing trend of non-local emission contribution is mainly due to ozone elevation in a larger scale. In the past few years, ozone levels in most of China were increasing, leading to continuous increase of non-local emission contribution over the PRD. This is added in lines 317-318 of the revised manuscript. The long-term trends of local emission contribution in different areas of the PRD is a combined results of VOCs and NO_x emission changes (as shown in Figure S1) and the ozone formation regime in particular sub-regions, which deserves further study.

8. There exist some uncommon usages of scientific writing English in the manuscript, such as Line 74 'such a philosophy' and Line 90 'is suffered from'.

Response: "such a philosophy" is changed as "this is" in line 76 of the revised manuscript. "Suffered from" is commonly used to describe the negative impacts from the thing afterwards.

9. In Line 127, u and v are wind direction and speed respectively while in Fig. 6 u and v both represent wind direction.

Response: In line 129, u and v refer to wind direction while m/s refers to wind speed. We add a "respectively" to make it clearer.

10. Some grammatical mistakes should be corrected, for example in Line 299, there should be a sentence after the word 'while'.

Response: We changed "while" to "and" in line 309 of the revised manuscript.

Quantitative impacts of meteorology and precursor emission changes on the long-term trend of ambient ozone over the Pearl River Delta, China and implications for ozone control strategy

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Abstract

China is experiencing increasingly serious ambient ozone pollution, including the economically developed Pearl River Delta (PRD) region. However, the underlying reasons for ozone increase remain
20 largely unclear, leading to perplexity in formulating effective ozone control strategies. In this study, by developing a statistical analysis framework combining meteorological adjustment and source apportionment, we examine quantitatively the impacts of meteorology and precursor emissions from within and outside the PRD on the evolution of ozone during the past decade. We found that meteorological condition has mitigated ozone increase, and its variation can account for at most 15% of
25 annual ozone concentration in the PRD. Precursor emission from outside the PRD ('non-local') makes the largest contribution to ambient ozone in the PRD and shows a consistently increasing trend, while that from within the PRD ('local') shows a significant spatial heterogeneity and plays a more important role during ozone episodes over southwestern. Under general conditions, the impact on northeastern is positive but decreasing, and on southwestern is negative but increasing. During ozone episodes, the
30 impact on northeastern is negative and decreasing, while on southwestern is positive but decreasing. Central and western PRD is the only area with increasing local ozone contribution. The spatial heterogeneity in both local ozone contribution and its trend under general conditions and ozone episodes are well interpreted by a conceptual model collectively taking into account ozone precursor emissions and their changing trends, ozone formation regimes, and the monsoonal and micro-scale synoptic conditions over different sub-regions of the PRD. In particular, we conclude that the inappropriate
35 NO_x/VOC control ratio within the PRD over the past years is most likely responsible for the ozone increase over southwestern, both under general conditions and during ozone episodes. By investigating the ozone evolution influenced by emission changes within and outside PRD during the past decade, this study highlights the importance of establishing a dichotomous ozone control strategy to tackle with
40 general conditions and pollution events separately. NO_x emission control should be further strengthened to alleviate peak ozone level during episodes. Detailed investigation is needed to retrieve appropriate NO_x/VOC ratios for different emission and meteorological conditions, so as to maximize the ozone reduction efficiency in the PRD.

45 **Keywords:** Ozone, Meteorological adjustment, Empirical orthogonal function, Ozone formation regime, Pearl River Delta

1. Introduction

Thanks to a series of stringent air pollution control measures, most types of air pollutants, including SO₂, NO_x, CO, PM₁₀ and PM_{2.5}, exhibited decreasing concentrations in the past six years (2013-2018) in China, with the only exception of ozone (Souri et al., 2017; Koukouli et al., 2018; Lin et al., 2018; Lu et al., 2018; Wang et al., 2018; Zhang et al., 2018). During 2015-2018, ozone concentrations in the three major city clusters, Beijing-Tianjin-Hebei, Yangtze River Delta, and Pearl River Delta (PRD), had increased by 20%, 4%, and 14%, respectively (Report on the State of the Environment in China, <http://english.mee.gov.cn/Resources/Reports/soe/>). Although with comparable median ozone concentrations, the magnitude and frequency of high-ozone events are much higher in China than those in Japan, South Korea, Europe, and the United States (Lu et al., 2018). Ozone would become one of the major air pollution control targets in China in the near future to protect public health.

Ozone control is far more difficult than particulate matter (PM) control, according to the experiences in Los Angeles and Mexico City (Madronich, 2014). The difficulties of ozone control lie in two major aspects. First, ozone can be contributed by both local formation and non-local transport, and their relative importance is largely driven by meteorological conditions and precursor emission characteristics (Elminir, 2005; Beaver and Palazoglu, 2009; Kovač-Andrić et al., 2009). Moreover, ozone is a secondary pollutant with non-linear relationship with its precursors, NO_x and volatile organic compounds (VOCs) (Stevenson et al., 2013; Thompson et al., 2014). Synergistic control with desirable VOC-to-NO_x reduction ratio is required for ozone reduction. However, such a ratio is hard to determine and practically implement due to our limited understanding on VOC emissions, especially those fugitive (Ou et al., 2016). The appropriate VOCs-to-NO_x reduction ratio may also vary greatly under different meteorological conditions. Therefore, from an ozone control point of view, it is fundamental to quantitatively understand the roles of meteorology and precursor emissions in shaping local and non-local ozone contributions, and their evolution during a long time scale in response to meteorology and emission changes.

Meteorology could either strengthen or dampen the efforts of precursor emission control on ozone reduction (Elminir, 2005; Beaver and Palazoglu, 2009; Kovač-Andrić et al., 2009). Hence, in order to investigate the effectiveness of precursor control during a long period, it is a common practice to homogenize meteorological conditions. In numerical simulation studies, ~~such a philosophy~~ is implemented by a set of scenarios in different meteorological and emission conditions (Gilliland et al., 2008; Godowitch et al., 2008; Wu et al., 2008; Foley et al., 2015). Differences in ozone levels between scenarios with the same meteorological conditions (emissions) are attributed to emission (meteorology) changes. Statistical models are also widely applied to establish relationship between ozone and meteorological variables so as to remove the meteorological impact, which is usually called meteorological adjustment (Lu and Chang, 2005; Zheng et al., 2007; Foley et al., 2015). After meteorological adjustment, ozone changes are solely attributed to emission changes.

Both local and non-local emission changes contribute to ambient ozone levels in a particular region. From an ozone control point of view, it is also essential to quantitatively differentiate local and non-local contributions. Source apportionment module coupled in chemical transport models, e.g. the Ozone Source Apportionment Technology (OSAT) in the Comprehensive Air-quality Model with extensions

(CAMx) and the Integrated Source Apportionment Method (ISAM) in the Community Multiscale Air Quality (CMAQ), are widely used to attribute ambient ozone concentrations at a particular place into different (local and non-local) source regions and categories (Li et al., 2013; Kwok et al., 2015). As numerical simulation is suffered from uncertainties in emission inventories and largely constrained in time span due to computing resources, statistical models, e.g. lowest-as-background method (Nielsen-Gammon et al., 2005; Xue et al., 2014) and Empirical Orthogonal Function (EOF) (Langford et al., 2009; Berlin et al., 2013), are preferentially adopted when the long-term monitoring data is available. They apportion local and non-local contributions by examining variability and co-variability of ozone concentrations at multiple monitoring sites. However, without meteorological adjustment, source apportionment by both methods reflects only the absolute contribution from local and non-local sources / processes and cannot directly link with local and non-local emissions of ozone precursors. Therefore, combined application of meteorological adjustment and source apportionment are indispensable in investigating the effect of local and non-local emission changes on long-term ambient ozone variations. Such combined application has not been reported in previous studies.

In this study, PRD is used as a research target area. After restraining its annual $PM_{2.5}$ concentration below $35\mu g m^{-3}$ (China's National Ambient Air Quality Standard for annual $PM_{2.5}$) for four consecutive years (2015-2018), PRD is the first major city cluster in China to transfer its main air pollution control target onto ozone. By utilizing continuous ozone monitoring at multiple stations across the PRD since 2007, we investigate the impacts of meteorology and local ('within PRD') and non-local ('outside PRD') emission changes on the long-term trend of ambient ozone by using the framework of meteorological adjustment followed by local and non-local contribution differentiation. Ozone contributions from meteorology and local and non-local emissions are quantitatively demonstrated in 2016 and 2017, the recent two years with significant ozone increase. We further develop a conceptual model depicting the impact of emission control within the PRD to the ambient ozone, both under general conditions and during ozone pollution episodes. Evaluation on the effectiveness of ozone precursor control measures within and outside the PRD during the past decade would shed light on future control efforts that hopefully shorten the ozone abatement paths experienced in Europe and the United States.

2. Data and Method

2.1 Ozone and meteorological data set

Hourly ozone monitoring data at fifteen monitoring stations across the PRD from 2007 to 2017 are used to calculate maximum daily 8-hour average (MDA8) in this study. Missing data are filled taking the yearly, monthly, weekly and hourly mean into account, otherwise it is replaced by the ozone data at the nearest monitoring station (Zheng et al., 2007). Geographical distribution of the monitoring stations is illustrated in Fig. 1, and the latitudes / longitudes and the types of functional areas where the stations are located are provided in Table 1.

The meteorological data during the same period, including daily maximum 2m temperature (T , °C), daily minimum relative humidity (RH, %), total net surface solar radiation (SSR, J/M^2), and 10m mean wind direction and speed (u and v , m/s, [respectively](#)), are retrieved from the European Center for Medium-range Weather Forecast (ECMWF) simulations for meteorological adjustment. Temporal and spatial resolution is 3-hour and $0.125^\circ \times 0.125^\circ$, respectively. Meteorological condition at the ozone monitoring station is represented by the simulation data at the closest point to the station, as illustrated in Fig. 1. In

this study, we composed an ozone and meteorological dataset with 4018 days at fifteen stations.

135 2.2 Meteorological adjustment

In this study, a statistical analysis framework combining meteorological adjustment and source apportionment is developed to identify ozone changes attributable to meteorology and local and non-local emissions. Long-term trends of ozone changes by meteorological conditions and local and non-local emissions are subsequently evaluated by trend analysis. In this study, ‘local’ emissions refer to those from within the PRD, while ‘non-local’ emissions refer to those from outside the PRD. A conceptual diagram highlighting major calculation procedures of the statistical analysis framework is shown in Fig. 2.

In meteorological adjustment, Kolmogorov-Zurbenko (KZ) filter is firstly used to separate the raw ozone and meteorological data into long-term, seasonal and short-term data (Rao and Zurbenko, 1994a; Rao and Zurbenko, 1994b). KZ filter can be expressed as

$$X(t) = LT(t) + SE(t) + ST(t) \quad (1)$$

Where $X(t)$ is the raw time series data, $LT(t)$ reflects the long-term trend in the time scale of years, $SE(t)$ is the seasonal variation in the time scale of months, and $ST(t)$ refers to short-term component in the time scale of days.

The KZ filter repeats the iterations of a moving average to remove the high-pass signal defined by

$$Y_i = \frac{1}{m} \sum_{j=-k}^k A_{i+j} \quad (2)$$

where k is the number of values included on each side, the window length $m=2k+1$, i is interval time, j is window variables, and Y is the input time-series. Thus the output of the i^{th} pass becomes the input for the $i+1^{\text{th}}$ pass, and so on. Different scales of motion are obtained by changing the window length and the number of the iterations (Milanchus et al., 1998; Eskridge et al., 1997). The filter periods of less than N days can be calculated with window length m and the number of iterations p , as

$$m \times p^{1/2} \leq N \quad (3)$$

So a $KZ(15, 5)$ filter with the window length of 15 with 5 iterations remove cycles of 33 days. The ozone and meteorological time series by $KZ(15, 5)$ refer to their baseline variations which are the sum of long-term $LT(t)$ and seasonal components $SE(t)$.

$$BL(t) = KZ_{(15,5)} = LT(t) + SE(t) = KZ_{(36, 53)} + SE(t) \quad (4)$$

The long-term trend is separated from the raw data by $KZ(365, 3)$ with the periods $>632d$, and then the seasonal and the short-term component $ST(t)$ can be derived by

$$SE(t) = KZ_{(15,5)} - KZ_{(36, 53)} \quad (5)$$

$$ST(t) = X(t) - BL(t) = X(t) - KZ_{(15,5)} \quad (6)$$

After KZ filtering, meteorological adjustment is conducted by stepwise regression between ozone concentration and meteorological factors such as T, RH and SSR (Flaum et al., 1996; Wise and Comrie, 2005; Papanastasiou et al., 2012).

$$A_{BL}(t) = a_{BL} + \sum b_{BLi} \cdot M_{BLi} + \epsilon_{BL}(t) \quad (7)$$

$$A_{ST}(t) = a_{ST} + \sum b_{STi} \cdot M_{STi} + \epsilon_{ST}(t) \quad (8)$$

$$\epsilon(t) = \epsilon_{BL}(t) + \epsilon_{ST}(t) \quad (9)$$

$$A_{ad}(t) = \epsilon(t) + \sum b_{BLi} \cdot \bar{M}_{BLi} + \sum b_{STi} \cdot \bar{M}_{STi} + a_{BL} + a_{ST} \quad (10)$$

Formula 7 and 8 are the multivariate regression models between baseline and short-term ozone and meteorological factors, respectively. $A_{BL}(t)$ and $A_{ST}(t)$ are the baseline and short-term components of

ozone and M_{BL} and M_{ST} are the baseline and short-term components of meteorological factors. The parameters a and b are the fitted parameters and i is time points (days). $\epsilon(t)$ is the residual term. The average meteorological condition \bar{M} of the same calendar date throughout 11 years is used as the base condition for that date, and the meteorological adjustment is conducted against the base condition. By doing so, the inter-annual variation of meteorology is removed while the annual variation is largely reserved. With the homogenized annual variation of meteorological conditions, $A_{ad}(t)$ in formula 10 represents the meteorologically adjusted ozone variations, and the difference between $X(t)$ and $A_{ad}(t)$ reflects the meteorological impact. It is noted that, by using the average meteorological condition as the base condition, the average ozone concentration during the 11 years keeps unchanged.

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2.3 Source apportionment of ozone contributions from local and non-local emissions

In this study, EOF and absolute principal component scores (APCS) are applied to apportion meteorologically adjusted ozone concentration into local and non-local emission sources. EOF transforms a large number of variables into a new set of uncorrelated, orthogonal principal components (PCs). The few new variables contains the most information of the original variables, and the new variables represent different processes contributing to ambient ozone levels. Here we present a brief description of EOF and APCS. Detailed information regarding the method can be found in Langford et al. (2009) and Berlin et al. (2013).

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EOF analysis is performed on the correlation matrix from the meteorologically adjusted ozone data set (4018 days \times 15 stations), without further rotation of the PCs. The first step is to normalize the ozone data (Thurston and Spengler, 1985; Guo et al., 2004).

$$Z_{ik} = (C_{ik} - C_i)/S_i \quad (11)$$

where C_{ik} is the concentration of ozone in sample k of the station i , C_i is the arithmetic mean value of ozone in station i and S_i is the standard deviation.

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$$Z_{ik} = L_{ip} \cdot P_{pk} \quad (12)$$

L_{ip} is loadings of EOF without rotation and P_{pk} is scores.

Since the factor scores are normalized with the mean to be zero, true zero is calculated through introducing an artificial sample with the zero concentration. Then the APCS are estimated by subtracting the artificial sample from the true samples.

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$$(Z_0)_i = \frac{(0 - C_i)}{S_i} = -C_i/S_i \quad (13)$$

$$(APCS)_{pi} = P_{pi} - P_{0pi} \quad (14)$$

The regression between APCS and ozone concentration estimates source contributions to C_i by

$$C_i = (b_0)_i + \sum APCS_p \times b_{pi} \quad (15)$$

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where $(b_0)_i$ is the constant term at station i , b_{pi} is the coefficient of the source p , and $\sum APCS_p$ is the scaled value of the factor p . Multiplication of $\sum APCS_p$ and b_{pi} calculates the contribution from source p to ozone concentration. Local and non-local sources are determined according to the temporal and spatial distribution characteristics of source contributions across the PRD.

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3. Results and Discussion

3.1 Long-term trend of meteorological impact on ozone concentration

Fig. 3a shows the long-term trends of ambient ozone, meteorologically adjusted ozone, and the meteorological impact in the PRD during 2007-2017. Ambient ozone concentration in the PRD increased

from $76 \mu\text{g m}^{-3}$ in 2007 to $89 \mu\text{g m}^{-3}$ in 2017, corresponding to an annual increase rate of $1.2 \mu\text{g m}^{-3}$.

220 Previous studies also evidenced ozone increase in the PRD (e.g. Li et al., 2014) and we here demonstrate that such an increase has been continuing for more than a decade. After meteorological adjustment, ozone concentration increases from $68 \mu\text{g m}^{-3}$ in 2007 to $90 \mu\text{g m}^{-3}$ in 2017, corresponding to an annual increase rate of $2.0 \mu\text{g m}^{-3}$. Higher increase rate of meteorologically adjusted ozone implies that if the meteorological condition keeps unchanged throughout the 11 years, ambient ozone concentration would
225 increase more significantly. As shown in Fig. 3a, meteorological conditions are generally favorable for ozone pollution during 2007-2011, responsible for at most $6 \mu\text{g m}^{-3}$ of ozone increase. During 2012-2017, meteorological condition became unfavorable for ozone pollution, leading to at most $6 \mu\text{g m}^{-3}$ of ozone reduction. Comparing between the most favorable (2007) and unfavorable (2016) year, meteorological condition change ozone concentration by $12 \mu\text{g m}^{-3}$ at most, roughly corresponding to 15% of annual ozone concentration. It is noted that the adjusted ozone concentration changed a little during 2011-2015. According to the VOCs and NO_x emissions in the PRD (Fig. S1), increase in VOCs emission started to mitigate in 2011, while NO_x emissions showed significant reduction starting from 2013. As PRD is generally in a VOC-limited ozone formation regime, reduction in the magnitude of VOCs emission increase is likely responsible for the minor changes in ozone during 2011-2015.
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235 It should be noted that meteorological adjustment does not change the overall increasing trend of ozone concentration, indicating that emission change is the primary driving factor for the long-term ozone trend. However, as shown in Fig. 3a, the fluctuation of ozone concentration is suppressed by meteorological adjustment, indicating that meteorology plays an important role in the inter-annual fluctuation of ozone
240 concentration, especially during 2011-2015 when ozone changes due to emissions are minor. During some specific period, meteorology plays a greater role in governing ozone changes than emissions, such as during 2016-2017 as will be discussed in section 3.4. As shown in Fig. 3b, meteorological adjustment significantly reduces the magnitude of ozone spikes, indicating that meteorological condition is one of
the most important driving factor for ozone episodes in the PRD.

245 Fig. 3c shows the impacts from different meteorological factors (T, RH, SSR, u and v) on ozone concentration. Overall, SSR is the most crucial factor and their variation follows well with that of the total meteorological impact. Contribution from the other four factors are comparable and relatively insignificant. As expected, higher SSR, higher T and lower RH are favorable for ozone production.

250 3.2 Spatial distribution of meteorological impact

We further examine the spatial distribution of meteorological impact. Fig. 4a shows the spatial distribution of averaged ozone concentration in the PRD and the annual ozone concentration changes before and after meteorological adjustment. Although northeastern PRD has the overall highest ozone
255 concentrations, central and western PRD shows the most rapid ozone increase during the 11 years (black bars), and such increases are further substantiated if meteorological impact is removed (green bars). There are two sub-regions in the PRD with overall decreased ozone concentrations, one in the northeast (TH and JGW) and the other in the southwest (ZML and TJ). The ozone decrease is largely mitigated or at ZML even reversed after meteorological adjustment. The different mechanisms leading to the ozone
260 increase in these two sub-regions are explained by a conceptual model in section 3.6.

The spatial distribution of meteorological impact in each year during 2007-2017 is illustrated in Fig. 4b.

It is noted that when the meteorological condition favors ozone pollution in the PRD, it increases more in the central and western area. On the contrary, when it decrease ozone concentration in the PRD, central and western PRD is also the region with larger decrease in most years. Therefore, central and western PRD is a meteorology-sensitive region for ozone pollution. [Generally, area with higher pollutant emission is more sensitive to changes in meteorological condition \(Seo et al., 2014\). Various studies have shown that central and western PRD is the area with the most intense VOCs and NO_x emissions over the PRD \(e.g. Zheng et al., 2009a\), therefore is more sensitive to meteorological condition.](#) Formulation of ozone control strategy in this region needs to consider meteorological impact.

3.3 Identification of ozone changes resulted from local and non-local emissions

Long-term variation of meteorologically adjusted ozone reflects the impacts from precursor (VOCs and NO_x) emission changes. As ozone can be contributed by both local production and long-range transport, it is important to quantitatively separate them from an emission control point of view. Considering PRD's monsoonal synoptic condition and that most local emissions are concentrated in the central [and western](#) PRD area, local ('within PRD') emissions tend to pose ~~distinct-contrasting~~ impacts to different sub-regions in different seasons while the impacts from non-local ('outside PRD') emissions in a relatively larger scale tend to be spatially ~~eonsistent-similar~~ over the PRD. We use this philosophy to examine the PCs derived from EOF analysis.

According to the Kaiser's rule (Wilks, 2006), three PCs are retained in EOF analysis, explaining 53%, 16% and 7% of total variance, respectively. Fig. 5a shows the interpolated PC loadings in the PRD, and Fig. 5b shows the long-term variation of PC scores during the 11 years. PC1 shows relatively consistent spatial distribution across the PRD, with its loadings ranging from 0.47 at TH to 0.84 at JJZ. Further examination on the relationship between PC scores and wind direction discovered that the score of PC1 is higher during high ozone concentration in the PRD, and is associated with northeasterly wind (Fig. 6a). Situated along the southeastern coast of China, PRD has two main prevailing winds, northeasterly mainly during winter and spring and southwesterly mainly during summer and fall. Northeasterly wind tends to bring emissions from inland to the PRD, while southwesterly wind originated from the ocean is relatively clean. All the above evidences support the notion that PC1 is associated with non-local impact from continental long-range transport. Higher impact in the central PRD may be caused by the rough land use and micro-scale circulation in this urbanized region that increases the residence time of non-local ozone. The score of PC1 (Fig. 5b) is almost consistently increasing during the 11 years, indicating increased ozone contribution from long-range transport.

In comparison, PC2 and PC3 loadings show significant spatial variations. PC2 loadings have an obvious north-south gradient with different signs, indicating that the impact of PC2 on northern and southern PRD are reversed at all times. Further examination on their relationship with wind direction, as shown in Fig. 6b and 6c, indicates that during high ozone periods, PC2 score tends to be positive with southerly winds and negative with northerly winds. With southerly winds, northern PRD receives the highest impact from PC2, leading to increased ozone concentration. On the contrary, southern PRD receives the highest impact (negative score and negative loading) with northerly winds. This reflects exactly the impact from emissions within the PRD posed by the north-south components of the prevailing winds. Similarly, PC3 is associated with the impact from local emissions by the west-east components of the prevailing winds. Therefore, PC2 and PC3 collectively reflect the impact of local emissions on ozone

formation. PC2 and PC3 scores show a bimodal pattern that are higher in 2007 and 2011-2014 (Fig. 5b). This suggests that local emissions pose higher ozone contribution to northern and eastern PRD during 2007 and 2011-2014 [while-and](#) to southern and western PRD during 2008-2010 and 2015-2017.

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The PC loadings and scores may reflect the spatial distributions and temporal variations of the PCs, respectively. However, as they are normalized values, APCS calculation is conducted to quantify the absolute ozone contributions from local and non-local emission sources. Fig. 7 shows ozone contributions from local and non-local emission sources at three representative stations, TH in northern, LH in central and DH in southern PRD, [while those at other stations are provided in Fig. S2](#). A first glance on this figure reveals identical and consistently increasing non-local trends at all three stations. [Actually, ozone level in most areas of China was increasing during the past years, which inevitably led to increased non-local contribution to ozone over the PRD \(Lu et al., 2018\)](#). Non-local emission contributions at DH reach 90~115 $\mu\text{g m}^{-3}$, more than doubling those of 44~56 $\mu\text{g m}^{-3}$ at TH. As explained previously, such a spatial heterogeneity is mainly caused by longer residence time of non-local ozone in the urbanized area. In comparison, local emission contributions show differences in both magnitude and trend over three stations. Local emission contribution on ozone ranges from 15~30 $\mu\text{g m}^{-3}$ at TH, 1~6 $\mu\text{g m}^{-3}$ at LH, and -30~-15 $\mu\text{g m}^{-3}$ at DH. As a net effect of ozone production and loss, the positive or negative sign of local emission contribution reflects the relative strengths of ozone production by HO_x and RO_x cycles and ozone loss by NO titration and deposition. As non-local emission contributions dominate over the local counterpart at all stations, its consistently increasing trend determines the meteorologically adjusted ozone trend over the 11 years.

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We further plot the spatial distribution of ozone contribution from local and non-local emissions and its long-term changes over the PRD, as shown in Fig. 8. Local emissions give positive contribution to northeastern, with the largest contribution of 31 $\mu\text{g m}^{-3}$ at JGW, and negative contribution to southwestern, with the largest contribution of -23 $\mu\text{g m}^{-3}$ at DH. Furthermore, apart from reversed ozone contribution from local emissions, northeastern and southwestern PRD also exhibit reversed trends in changes of ozone contribution from local emissions during the 11 years, as illustrated in the bars of Fig. 8. The most significant increase trend is found over southwestern, with the largest increase rate of 0.6 $\mu\text{g m}^{-3} \text{ year}^{-1}$ at DH, while the most significant decrease trend is found over northeastern, with the largest decrease rate of 0.8 $\mu\text{g m}^{-3} \text{ year}^{-1}$ at JGW. The underlying mechanism resulting in the opposite trends in both local ozone contribution and its long-term changes between northeastern and southwestern are explained with a conceptual model in section 3.6. In comparison, the ozone contributions from non-local emissions are relatively consistent over the PRD, and non-local emission poses increasing influence on ozone for the entire region.

3.4 Identification of driving factors for ozone changes in 2016 and 2017

With meteorological adjustment and source apportionment, the contributions from meteorology and local and non-local emissions to the ambient ozone changes can be quantitatively analyzed for all years at all stations. In this section, we demonstrate this capability by analyzing the relative importance of meteorology and local and non-local emissions to the ozone changes in the recent two years, 2016 and 2017. Significant ozone level increases are revealed at most stations during the two years, with the average [annual](#) concentration rising from 81 $\mu\text{g m}^{-3}$ in 2016 to 87 $\mu\text{g m}^{-3}$ in 2017. It is found that meteorology, local emission and non-local emission contribute to around 3.5 $\mu\text{g m}^{-3}$, -0.1 $\mu\text{g m}^{-3}$ and 2.0

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$\mu\text{g m}^{-3}$ of ozone increase, respectively. Overall, meteorology plays a greater role in elevating ozone levels during these two years.

Contributions from meteorology and local and non-local emissions are further analyzed at each monitoring station, as listed in Table 2. Under general conditions, in comparison with local and non-local emissions, meteorology gives the highest contributions to ozone changes at all stations except for CZ and DH, the two southwestern-most stations. In addition, local emissions gives higher contributions than non-local ones at CZ, DH, JJZ, ZML and TJ, the cluster of stations in the southwestern PRD. Therefore, the ozone increase over southwestern PRD during these two years is most attributable to local emission changes, while the ozone increase in other parts of the PRD is firstly driven by meteorological condition changes, followed by non-local emission changes. This suggests that in order to reduce ozone levels in the southwestern PRD, strengthening local VOCs emission control should be of the top priority, so as to prevent ozone titration from decreasing further.

3.5 Impact of meteorology and emission changes during ozone episodes

In this section, we examine the impacts of meteorology and local and non-local emission changes to ambient ozone level during ozone episodes. Ozone episodes are defined as days with MDA8 ozone concentration greater than $160 \mu\text{g m}^{-3}$ at five stations or more across the PRD. As shown in Fig. S4S3, ozone levels are the highest in the central PRD, mainly Guangzhou and Foshan, during ozone episodes.

Fig. S2-S4 shows the long-term trends of ambient ozone, meteorologically adjusted ozone, and the meteorological impact in the PRD during ozone episodes in 2007-2017. Ambient ozone concentration during episodes increases from $150 \mu\text{g m}^{-3}$ in 2007 to $161 \mu\text{g m}^{-3}$ in 2017, corresponding to an annual increase rate of $1.0 \mu\text{g m}^{-3}$. It is noteworthy that meteorological adjustment does not alter ozone concentration much, with the largest change of $3 \mu\text{g m}^{-3}$ only. This implies that, although with significant variation under general conditions, meteorology does not vary significantly during ozone episodes across all years. Changes in precursor emissions are therefore the driving factor for long-term ozone variations during ozone episodes. A slightly different picture is discovered in 2017 during which meteorology is the major culprit for ozone increase. Without meteorological impact, ozone level during episodes should be lower than that in 2016.

We further differentiate ozone changes into those by local and non-local emissions using EOF/APCS approach. Four principal components are discovered, and they are assigned to local or non-local emissions by their spatial variations, as shown in Fig. S3S5. Fig. 9 illustrates the long-term trend of ozone contribution by local and non-local emissions during ozone episodes at TH, LH and DH stations. At TH and LH, non-local emissions give dominant contribution to ozone, while local emissions pose negative impacts, while contributions from local and non-local emissions are comparable at DH. Different from general conditions during which non-local contribution shows a consistently increasing trend, non-local contribution fluctuates greatly during ozone episodes and presents a bimodal picture. Starting from 2014, ozone contribution from non-local emissions has leveled off and decreased gradually. [This may be related to the VOCs and NO_x emissions in the non-PRD area in Guangdong \(Fig. S1\) and further upwind area, which deserves further study.](#)

Local emission contribution to ozone during episodes differs greatly in different areas. As shown in Fig.

395 10a, local emissions give positive contribution to southwestern, with the largest average contribution of
78 $\mu\text{g m}^{-3}$ at DH. They pose negative contribution to northeastern, with the largest contribution of -36 μg
400 m^{-3} at TH and HG. Such a spatial distribution is contrary to that during general conditions, as illustrated
in Fig. 8a. Stations over central PRD show increasing trend, with the largest increase rate of 1.9 $\mu\text{g m}^{-3}$
year⁻¹ at HG, while stations surrounding central and western PRD show decreasing trend, while the
largest decrease rate of 3.5 $\mu\text{g m}^{-3}$ year⁻¹ at JGW. In comparison, ozone contributions from non-local
emissions are relatively consistent over the PRD, with the hotspot shifted from central and western PRD
under general conditions (Fig. 8b) to central and eastern PRD during ozone episodes (Fig. 10b). The
entire PRD experienced increasing ozone contribution from non-local emission. Comparing with non-
405 local emission, ozone contribution from local emission and its trend show significant spatial
heterogeneity. We develop a conceptual model to explain in detail the underlying mechanisms resulting
in the distinct spatial distribution of local ozone contribution and its trend between general conditions
and ozone episodes, as elaborated in section 3.6.

3.6 A conceptual model describing impact of local emission changes to ozone in the Pearl River

410 Delta

As discussed in section 3.3 and 3.5, the spatial pattern of ozone contribution from local ('within PRD')
emissions and its long-term changes in the PRD under general condition and during ozone episodes
present different pictures. Under general condition, local emissions give positive and decreasing
contribution to ozone over northeastern PRD, and negative and increasing contribution over southwestern
415 (Fig. 8a). In contrast, during ozone episodes, local emissions give negative and decreasing contribution
over northeastern, and positive and decreasing contribution over southwestern. Central and western PRD
is the only region having slight increasing local ozone contribution during episodes (Fig. 10a). In this
section, we aim to provide detailed explanation on such phenomena by developing a conceptual model
collectively taking into account ozone precursor emissions and their changing trends, ozone formation
420 regimes, and the monsoonal and micro-scale synoptic conditions over different sub-regions of the PRD.

3.6.1 General condition

PRD has distinct VOCs and NO_x emission characteristics across its different sub-regions, leading to
different prevailing ozone formation regimes (OFR) over the PRD. Central PRD, essentially western and
425 southern Guangzhou, Foshan and western Dongguan, is the area with the most significant economic and
industrial activities. Central PRD is associated with significant amount of VOCs and NO_x emissions
(Zheng et al., 2009b, Zhong et al., 2018), and is mostly in a VOC-limited OFR (Ye et al., 2016). The
polluted air mass can be transported to different areas of the PRD under different prevailing winds, and
largely determines the ozone behaviors over those areas. In the past years, NO_x emissions are decreasing
430 due to stringent control measures, while VOCs emissions are increasing, as shown in Fig. S4S1.

Northeastern PRD is mainly a rural area with plenty of vegetation coverage. Significant VOC emissions
from biogenic sources make it primarily in a NO_x-limited OFR, especially in summer (Ye et al., 2016).
In summer and fall, southwesterly winds originated from the South China Sea prevail, bringing the NO_x-
435 laden air mass from central PRD to the downwind NO_x-limited northeastern and increasing ozone levels
over TH, XP and JGW stations. However, NO_x/VOC ratio in the air mass is decreasing during the past
years due to emission control measures that are preferentially targeting on NO_x emissions. Lowered
NO_x/VOC ratio would inhibit ozone production in the NO_x-limited northeastern, leading to a downward

ozone trend. In contrast, southwestern PRD shows relatively higher NO_x/VOC emission ratios, and is
440 mostly in a VOC-limited OFR (Ye et al., 2016). The OFR would shift to be more VOC-limited in winter
due to the suppressed biogenic VOC emissions and reduced reaction rate of HO_x and RO_x cycles. In
winter and spring, northeasterly winds originated from the Eurasia Continent prevail, bringing the NO_x -
laden air mass from the central PRD to the southwestern. The NO_x -laden air mass would react
preferentially with ozone in the VOC-limited southwestern, thereby decreasing the ozone levels at CZ,
445 DH and ZML stations. Due to the strengthened NO_x emission control that reduces NO_x/VOC ratio from
the central PRD, ozone titration is largely mitigated, leading to an upward ozone trend over southwestern
in the past few years. Fig. 11 provides a conceptual diagram on the impact of local emission control on
ozone concentrations and their changing trends over the PRD.

450 Hence, the combined influences by reduced ozone titration from local emissions and increased ozone
import from non-local emissions make southwestern the area having the most rapid ozone increase over
the PRD. In order to curb ozone increase in the southwestern, VOC emission control within the PRD
must be strengthened to elevate NO_x/VOC ratio into a level that ozone titration would not be further
reduced. With decreased influence from local emissions, northeastern shows the least ozone increase.

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3.6.2 Ozone pollution episodes

Both meteorology and precursor emissions exhibit significant differences during ozone episodes in
comparison with general conditions. Ozone episodes typically happen in summer and fall with hot and
sunny weather and weakened background wind, which is very often associated with a high pressure ridge
460 or approaching of a tropical cyclone (Huang et al., 2006). Temperature very often rises above 32 degree
Celsius with abundant sunshine, leading to more intense biogenic VOC emissions over the PRD.
Considering NO_x emissions are insensitive to temperature rise and the high reactivity of biogenic VOCs,
the effective NO_x/VOC ratio becomes much lower. As a result, NO_x -limited OFR over northeastern is
intensified, and VOC-limited OFR over southwestern shifts to NO_x -limited. VOC-limited area shrinks to
465 merely central PRD and the magnitude is largely weakened (Wang et al., 2011; Jin and Holloway, 2015).
Due to significant NO_x emissions, the urban central PRD is probably the last area turning into NO_x -
limited due to enhanced biogenic VOC emissions during ozone episodes.

In addition, the prevailing wind direction changes from northeasterly / southwesterly to easterly, as shown
470 in Fig. S5S6. With weakened background wind, micro-scale circulations such as land-sea breeze develop
around the Pearl River Estuary (PRE), and becomes an effective mechanism in trapping and mixing up
pollutants emitted surrounding the PRE (Lo et al., 2006). Micro-scale circulations increase the residence
time of pollutants over the PRE and thus expedite chemical reactions to produce ozone. High ozone
produced around the PRE is brought to southwestern PRD (a 'sink' region) by the weak easterly wind,
475 thereby increasing ozone levels at DH, ZML and TJ stations. In contrast, with easterly wind, northeastern
receives little impact from the central PRD ozone hotspot while instead serves as a 'source' region (ozone
import from further east is accounted for as impact from non-local emissions), thereby providing negative
contribution at TH, XP, JGW, HG and LH stations.

480 With higher biogenic VOC emissions and VOC oxidation rate, the OFR distribution over the PRD during
ozone episodes vary from that under general conditions. The preferential NO_x emission reduction due to
stringent control would lead to downward trend of local ozone contribution over northeastern due to

intensified NO_x-limited OFR, also downward trend over southwestern due to shift from VOC-limited to NO_x-limited OFR. An upward trend is only discovered over central and western PRD (HG, LH, HJC and ZH) where NO_x emissions are very strong and still persist in VOC-limited OFR. Fig. 12 provides a conceptual diagram on the impact of local emission control on ozone concentrations and their changing trends over the PRD during ozone episodes.

Hence, even with different formation mechanisms from general conditions, southwestern PRD, mainly Zhongshan, Zhuhai and eastern Jiangmen, is still the area with the most significant impact from local emissions during ozone episodes. However, with less NO_x emissions than central PRD, OFR over southwestern has shifted from VOC-limited to NO_x-limited, leading to reduced local ozone contribution. Comparison of different trends between central and southwestern PRD actually highlights the fact that NO_x emission control is one of the possible means to reduce ozone levels over the PRD, especially during ozone episodes with significantly enhanced biogenic VOCs emissions. Further reduction of NO_x emissions, after bypassing the optimal effective NO_x/VOC ratio [leading to the highest ozone concentration](#), would rapidly pull down peak ozone level and eventually bring it into attainment (Ou et al., 2016). Different OFR characteristics under general condition and during ozone episodes also highlight the importance of formulating dynamic control measures tailored for different emission and meteorological conditions.

4. Conclusion and Implication

Ambient ozone level in a particular area is determined by the interaction between meteorology and emission of ozone precursors, VOCs and NO_x. Differentiation of their impacts are important to evaluate the effectiveness of emission control measures in the past and to shed light on directions for future control plans. In this study, we develop a statistical analysis framework to identify ozone changes attributable to meteorology and local and non-local emissions in the PRD. The framework is essentially a combination of meteorological adjustment and source apportionment by EOF. We found that meteorology does not alter the increasing trend of ozone during 2007-2017, but significantly mitigate the magnitude of increasing. Ozone increase solely due to precursor emission changes would have been more significant.

In comparison with non-local precursor emissions, the impacts of local precursor emissions on ambient ozone present significant spatial and temporal heterogeneity over the PRD. Northeastern and southwestern exhibit different net ozone production and loss characteristics under general conditions and during ozone episodes. In response to the preferential NO_x emission control during the past years, local ozone contribution decreases over northeastern and increases over southwestern under general conditions, while decreases over both northeastern and southwestern but increases over central and western PRD during ozone episodes. Such a complex characteristics can be well interpreted by a conceptual model collectively taking into account ozone precursor emissions and their changing trends, ozone formation regimes, and the monsoonal and micro-scale synoptic conditions over different sub-regions of the PRD. In particular, OFR shift during ozone episodes in response to higher biogenic VOC emissions and VOC oxidation rate is the fundamental cause for different trends both spatially and temporally. We conclude that the past control measures preferentially targeted on NO_x are most likely responsible for ozone increase in the PRD, especially over southwestern by reduced ozone titration. However, OFR has started to shift from VOC-limited to NO_x-limited over southwestern, especially during ozone episodes. Therefore, NO_x emission control should be further strengthened to alleviate peak ozone levels.

By investigating the ozone evolution influenced by emission changes within and outside PRD during the past decade, this study highlights the complexity in ozone pollution control in the PRD. The complexity lies in three aspects. First, ozone control should be location-specific. Northeastern is the area benefited from current control measures in the PRD, and the main focus should be on co-prevention and co-control with further northeastern areas, e.g. Jiangxi and Fujian, to reduce long-range transport; Central and southwestern PRD should pay more efforts on VOCs control to elevate NO_x/VOC ratio into a level that ozone titration would not be further reduced. Second, ozone control should be temporally dynamic and largely dependent upon meteorological conditions. OFR may change greatly under different meteorological conditions which would influence effective control strategy and deserve more in-depth investigation. In particular, precursor emissions surrounding the PRE should be preferentially controlled during ozone episodes as they may contribute greatly to ozone formation when trapped over PRE by the micro-scale circulations. They are responsible for ozone hotspot over southwestern with a drastically increasing trend. Last but not least, under every circumstance, the most desirable NO_x/VOCs ratio for emission control should be investigated in detail. For example, control measures during ozone episodes should preferentially target on NO_x in the context of significantly enhanced biogenic VOCs emissions. Comparison of different trends between central and southwestern PRD provides a perfect highlight on the effect of NO_x control. Further reduction of NO_x emissions, after bypassing the optimal effective NO_x/VOC ratio, would rapidly pull down peak ozone level and eventually bring it into attainment (Ou et al., 2016).

Authorship Contribution Statement

ZY and JZ designed the experiments and LY, HL, ~~and~~ XL and YB carried them out. PKKL and DC provided ozone monitoring data and contributed to the discussion of the results. LY and ZY drafted the paper, with all co-authors contributing to subsequent enhancements.

*These authors contribute equally to this article.

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Table 1. Location of fifteen ozone monitoring stations across the Pearl River Delta and their environmental background.

Station	Full name	City	Longitude (E)	Latitude (N)	Environmental Background
CW	Central/Western	Hong Kong	114.15	22.28	Residential/Commercial
CZ	Chengzhong	Zhaoqing	112.47	23.05	Residential/Commercial
DH	Donghu	Jiangmen	113.08	22.59	Urban
HG	Haogang	Dongguan	113.73	23.03	Residential/Commercial
HJC	Huijingcheng	Foshan	113.10	23.00	Residential/Commercial
JGW	Jinguowan	Huizhou	114.38	22.93	Residential
JJZ	Jinjuzui	Foshan	113.26	22.81	Suburban
LH	Luhu	Guangzhou	113.28	23.15	Urban
LY	Liyuan	Shenzhen	114.09	22.55	Urban
TC	Tung Chung	Hong Kong	113.91	22.27	Residential
TH	Tianhu	Guangzhou	113.62	23.65	Rural
TJ	Tangjia	Zhuhai	113.58	22.34	Commercial/Industrial
XP	Xiapu	Huizhou	114.40	23.07	Commercial
YL	Yuen Long	Hong Kong	114.02	22.44	Residential
ZML	Zimaling	Zhongshan	113.40	22.50	Residential/Commercial

710 **Table 2. Contributions of meteorology and local and non-local emission changes to the ozone concentration change ($\mu\text{g m}^{-3}$) in 2016-2017 at fifteen monitoring stations in the Pearl River Delta under general conditions.**

Station	CZ	DH	JJZ	ZML	TJ	HJC	LH	LY
Meteorology	1.9	3.1	3.3	4.0	5.3	3.2	4.0	3.2
Local	3.5	4.8	2.9	3.4	2.2	2.2	-0.3	0.1
Non-local	2.2	2.7	2.8	2.6	2.0	2.5	2.3	1.5
Station	YL	TC	CW	HG	XP	JGW	TH	PRD
Meteorology	3.8	1.9	2.9	3.9	4.2	4.5	4.0	3.5
Local	0.4	0.3	-0.2	-2.7	-5.5	-6.3	-5.0	-0.1
Non-local	1.8	0.9	1.4	2.4	1.8	1.6	1.3	2.0

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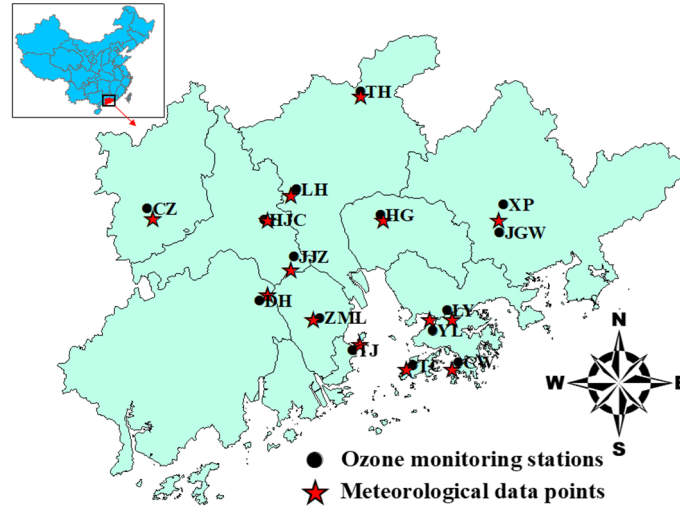
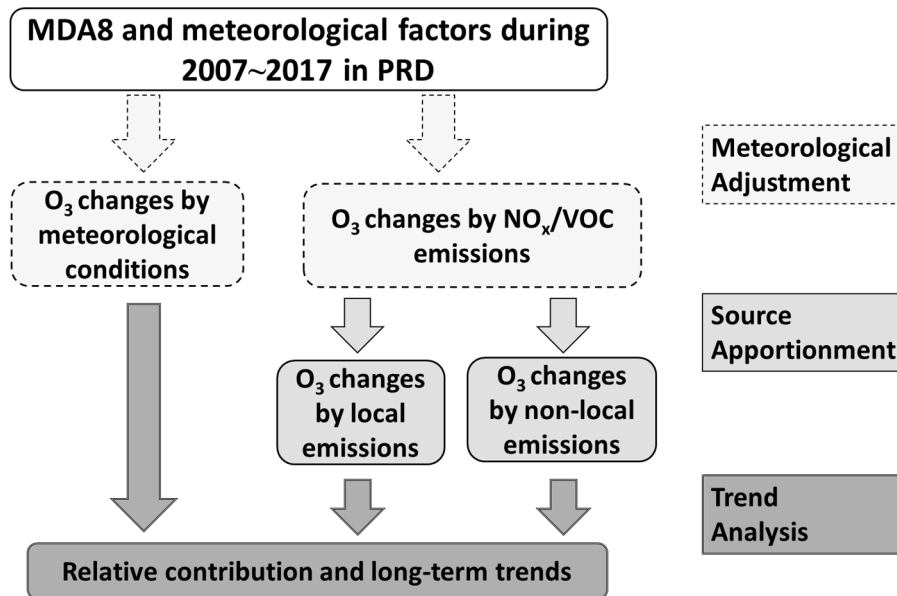
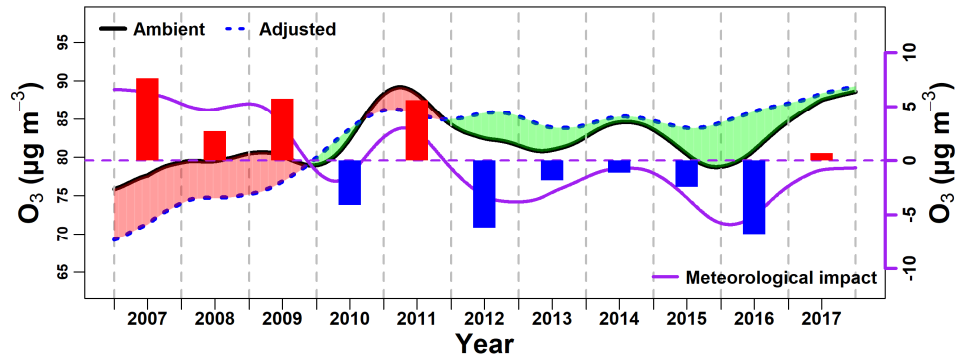


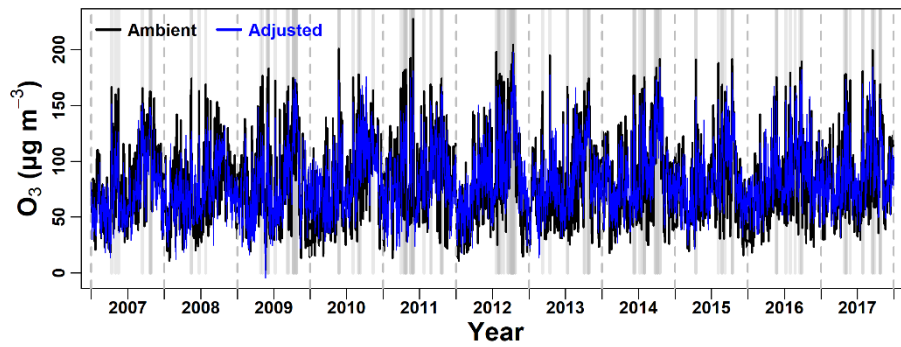
Fig 1. Distribution of ozone monitoring stations and meteorological data points in the Pearl River Delta.



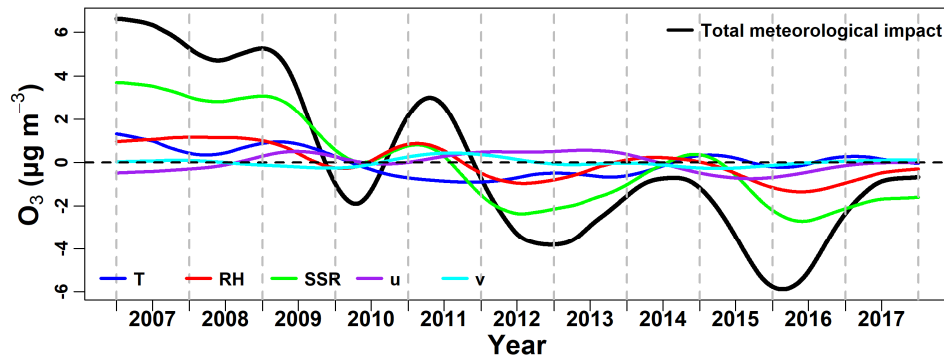
720 Fig 2. Flowchart of the statistical analysis framework to identify the long-term impacts of meteorology and local and non-local emissions on ambient ozone.



(a)



(b)



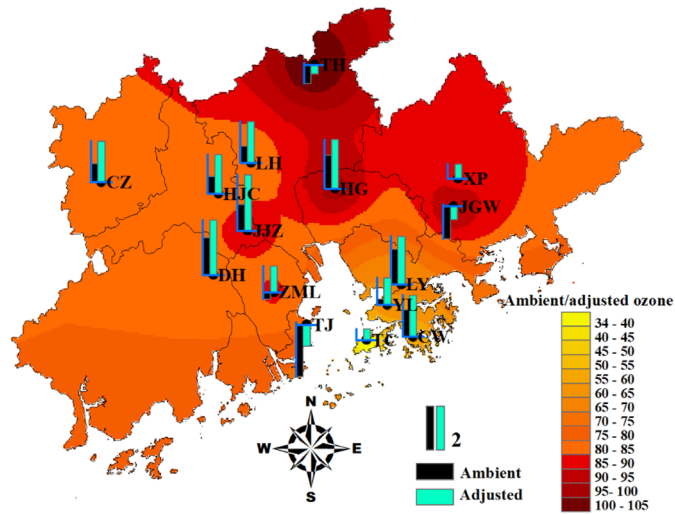
(c)

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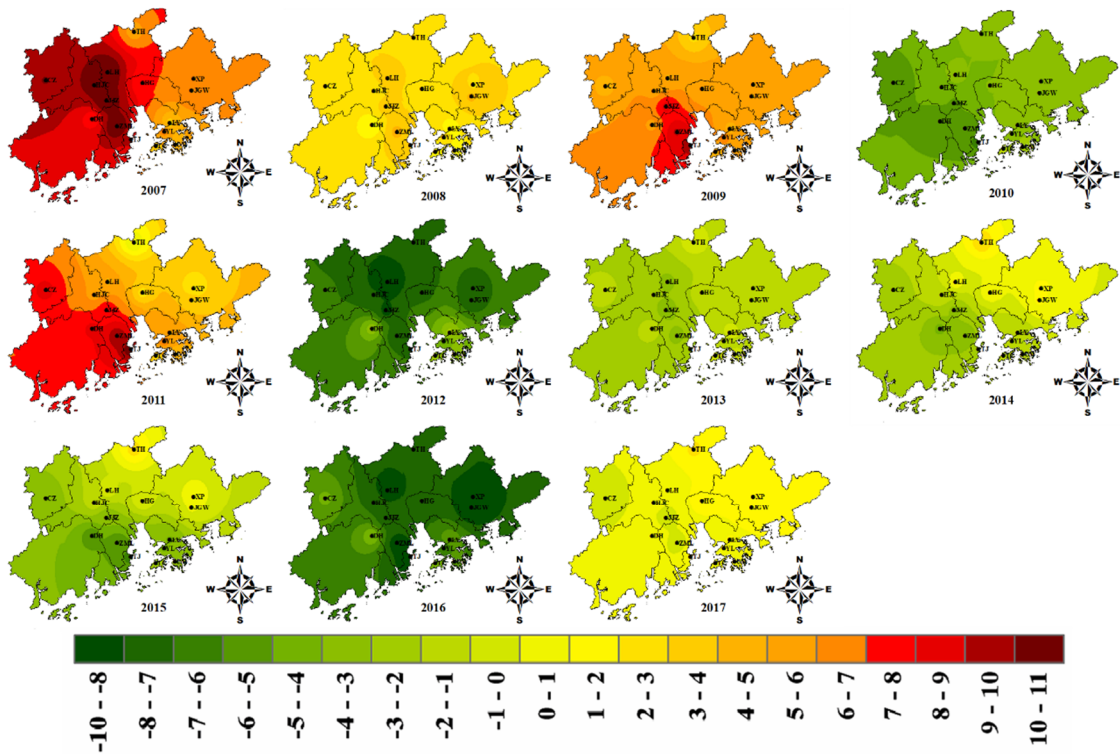
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Fig 3. (a) Long-term trends of ambient ozone, meteorologically adjusted ozone, and the meteorological impact in the Pearl River Delta during 2007-2017. Periods with positive and negative meteorological impacts are shadowed in red and green, respectively. Red and blue bars represent ozone increase and reduction attributed to meteorology in each year, respectively. (b) Ozone concentration time series before (black) and after (blue) meteorological adjustment. Gray areas represent periods with ozone concentration over $160 \mu\text{g m}^{-3}$. (c) Long-term variations of meteorological impact by different meteorological factors.

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(a)

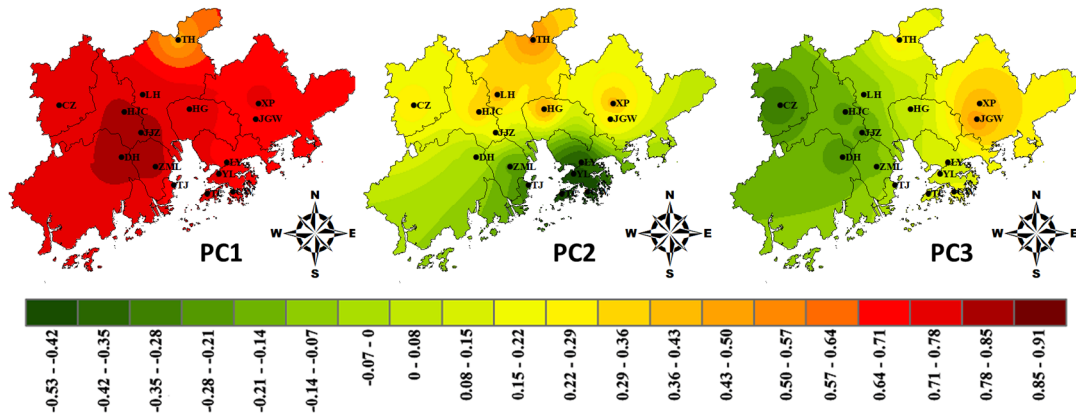


(b)

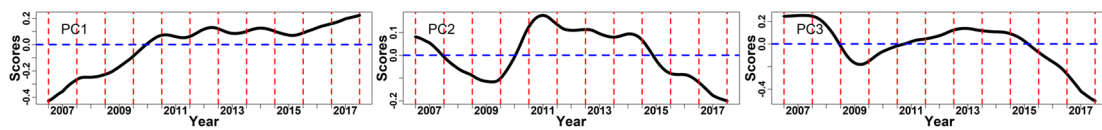
Fig 4. (a) Spatial distribution of averaged ozone concentrations ($\mu\text{g m}^{-3}$) in the Pearl River Delta and annual ozone changes ($\mu\text{g m}^{-3} \text{ year}^{-1}$) before and after meteorological adjustment over the fifteen monitoring stations during 2007-2017. The bar length in the legend corresponds to an annual increase of $2 \mu\text{g m}^{-3}$. (b) Annual variation of meteorological impact on ozone concentration ($\mu\text{g m}^{-3}$) over the Pearl River Delta during 2007-2017.

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(a)



(b)

Fig 5. (a) Spatial distribution of principal component loadings in the Pearl River Delta, and (b) long-term variation of principal component scores during 2007-2017. PC1 reflects non-local emission impacts while PC2 and PC3 refer to impacts from different local emissions.

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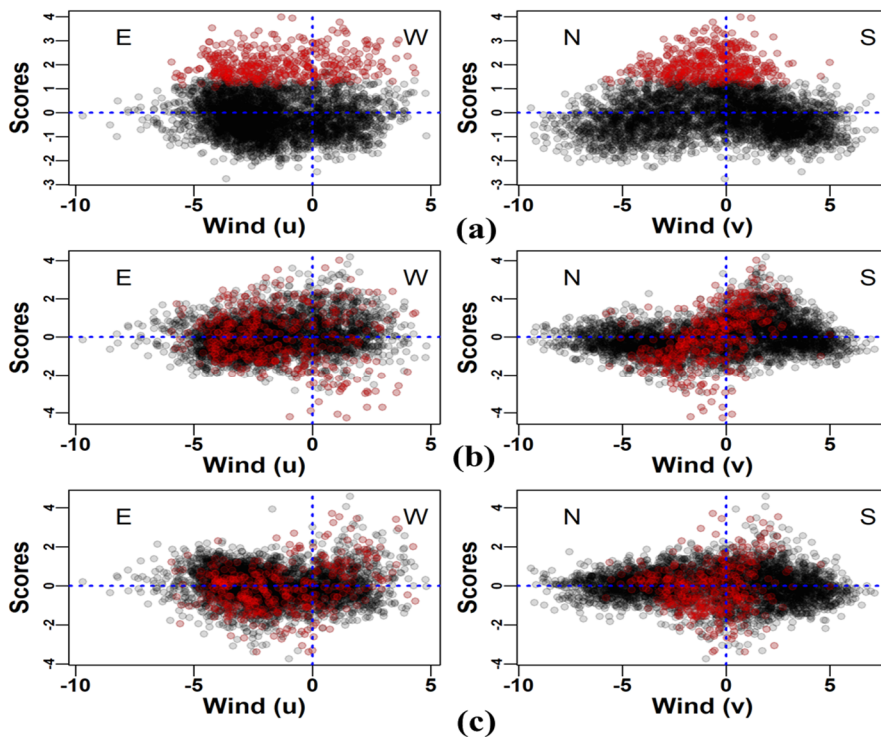
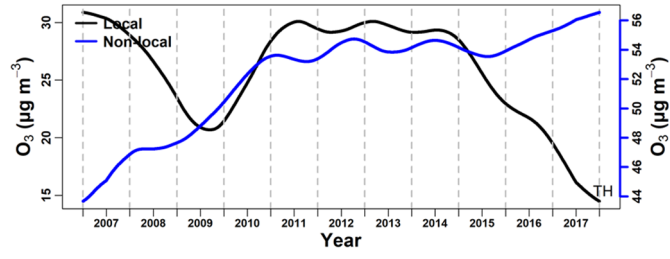
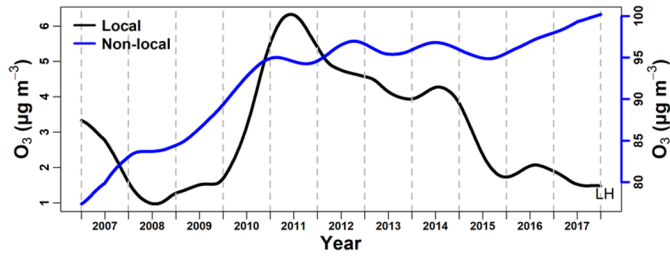


Fig 6. Scatterplot between principal component scores (a-c: PC1-3) and wind (u and v). u and v are the east-west and north-south components of wind (u: +west/-east, v: +south/-north). Red points refer to samples with high ozone concentration (over 90th percentile).

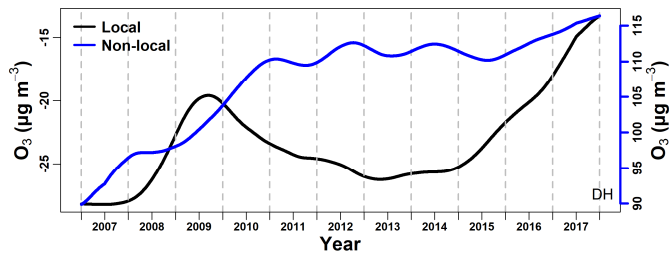
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(a)



(b)

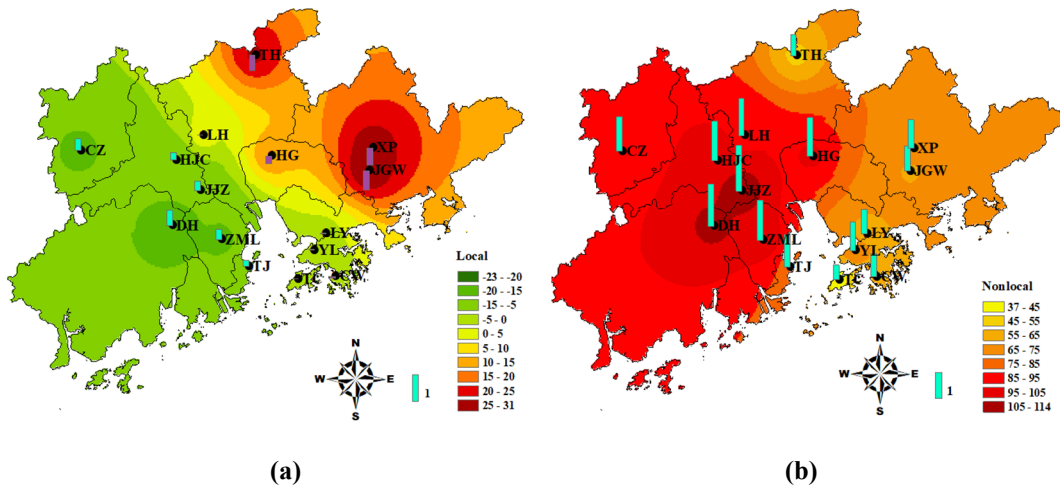


(c)

Fig 7. Long-term trend of ozone contributed by local (black) and non-local (blue) emission sources from 2007 to 2017 at (a) TH, (b) LH and (c) DH stations.

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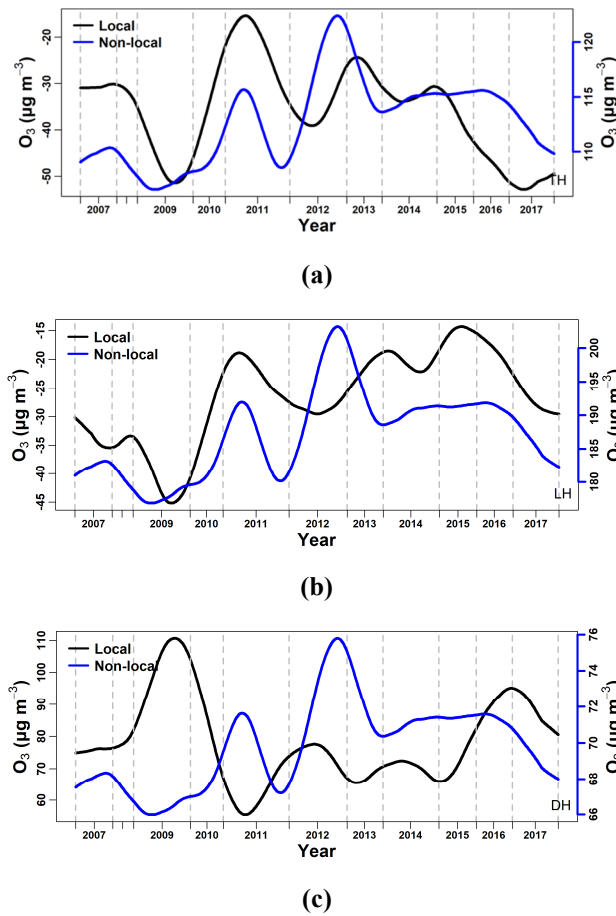
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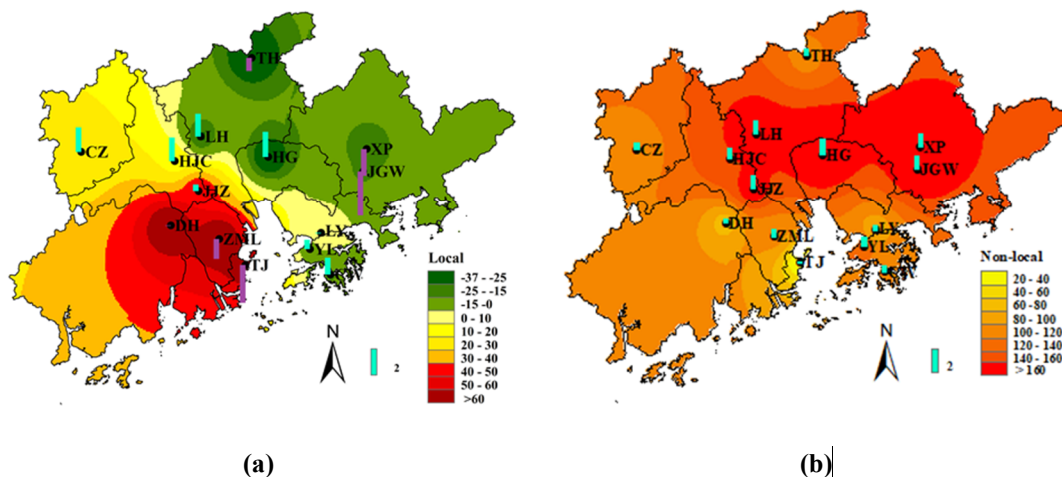
Fig 8. Spatial distribution of ozone contribution from (a) local and (b) non-local emissions ($\mu\text{g m}^{-3}$) of each station and their annual change rate in the Pearl River Delta. Bars in blue above / in purple below the station point indicate increasing / decreasing contributions. The bar length in the legend corresponds to an annual increase of $1 \mu\text{g m}^{-3}$. Ozone contributions from local emissions show positive but decreasing trend in the northeastern and negative but increasing trend in the southwestern. Ozone contributions from non-local emissions are positive and increasing region-wide.

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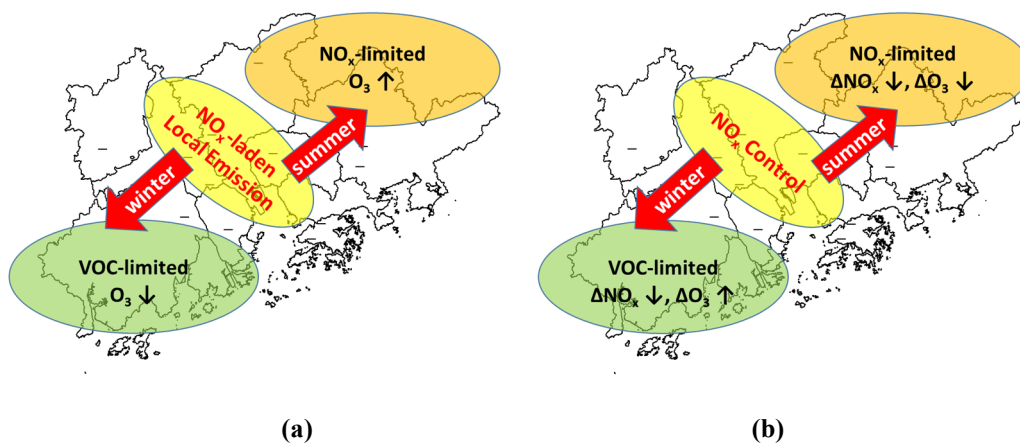
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Fig 9. Long-term trend of ozone contributed by local (black) and non-local (blue) emission sources from 2007 to 2017 at (a) TH, (b) LH and (c) DH stations during ozone episodes. Inver



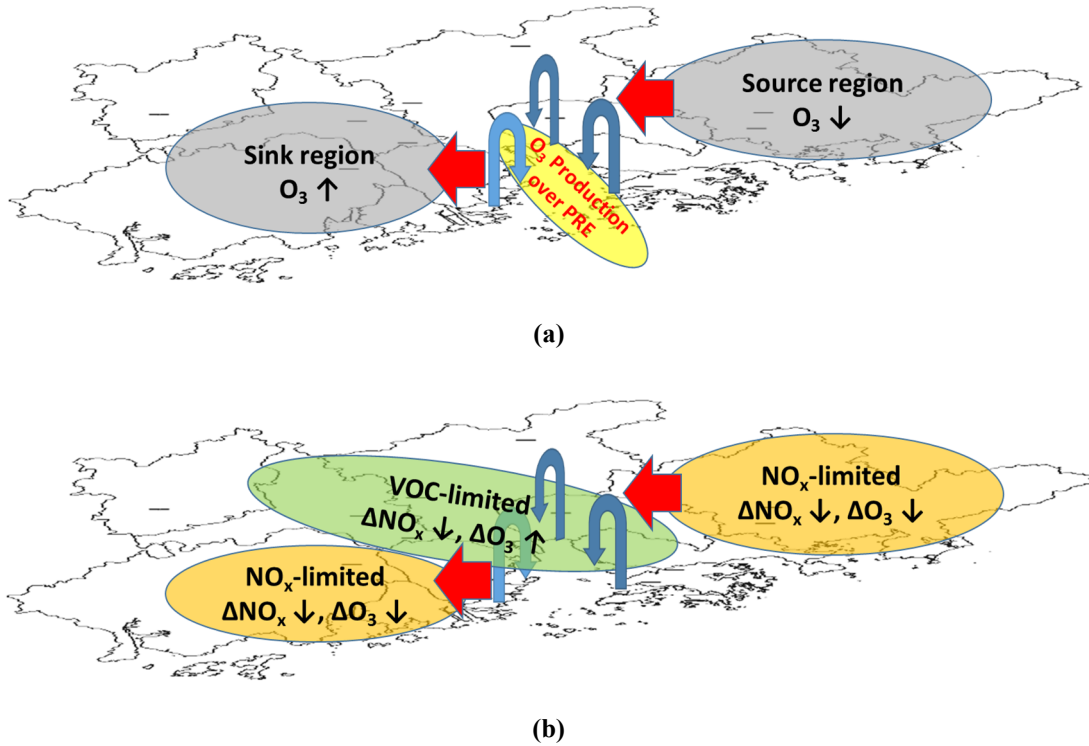
790 Fig 10. Spatial distribution of ozone contribution from (a) local and (b) non-local emissions ($\mu\text{g m}^{-3}$) of each station and their annual change rate in the Pearl River Delta during ozone episodes. Bars in blue above / in purple below the station point indicate increasing / decreasing contributions trend. The bar length in the legend corresponds to an annual increase of $2 \mu\text{g m}^{-3}$. Ozone contributions from local emissions are positive in the southwestern and negative in the northeastern. Central PRD is the only area with increasing local ozone contribution trend. Ozone contributions from non-local emissions are positive and increasing region-wide.

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800 Fig 11. A conceptual diagram on the impacts of local emissions on (a) ozone concentrations and (b) their changing trends over the Pearl River Delta. Local NO_x -laden emissions increase ozone level ($\text{O}_3 \uparrow$) in the downwind northeastern in summer and fall, but the increase is suppressed due to the preferential NO_x control ($\Delta \text{NO}_x \downarrow$), leading to net ozone decrease ($\Delta \text{O}_3 \downarrow$). In comparison, local emissions decrease ozone level ($\text{O}_3 \downarrow$) in the downwind southwestern in winter and spring, but the decrease is also mitigated due to NO_x control ($\Delta \text{NO}_x \downarrow$), leading to net ozone increase ($\Delta \text{O}_3 \uparrow$). Such a phenomenon is essentially governed by different ozone formation regimes in northeastern (NO_x -limited) and southwestern PRD (VOC-limited).

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810 Fig 12. A conceptual diagram on the impact of local emissions on (a) ozone concentrations and (b) their
 815 changing trends in the Pearl River Delta during ozone pollution episodes. The blue curved angles indicate
 micro-scale circulations such as land-sea breeze developed around the Pearl River Estuary (PRE). The
 micro-scale circulation leads to a high ozone area around the PRE due to effective mixing and reaction
 between VOCs and NO_x. High ozone is transported to southwestern by easterly wind, increasing local ozone
 contribution. In comparison, northeastern is a source region of ozone therefore its local contribution is
 negative. With higher biogenic VOC emissions and VOC oxidation rate during ozone episodes, most of PRD
 is in the NO_x-limited ozone formation regime except for urban central PRD which is still VOC-limited due
 to intense NO_x emissions. Therefore, reduced NO_x emissions ($\Delta \text{NO}_x \downarrow$) lead to decreasing ozone level (ΔO_3
 \downarrow) over both northeastern and southwestern and increasing ozone level over central PRD ($\Delta \text{O}_3 \uparrow$).

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Supplementary Material for

Quantitative impacts of meteorology and precursor emission changes on the long-term trend of ambient ozone over the Pearl River Delta, China and implications for ozone control strategy

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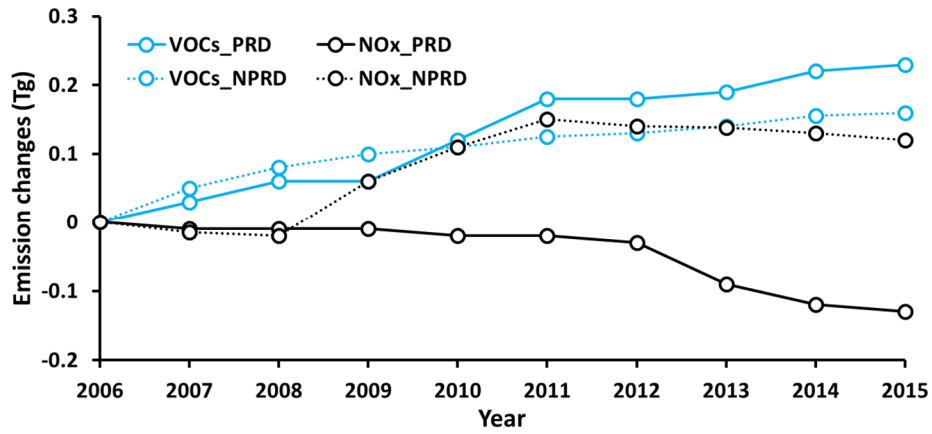
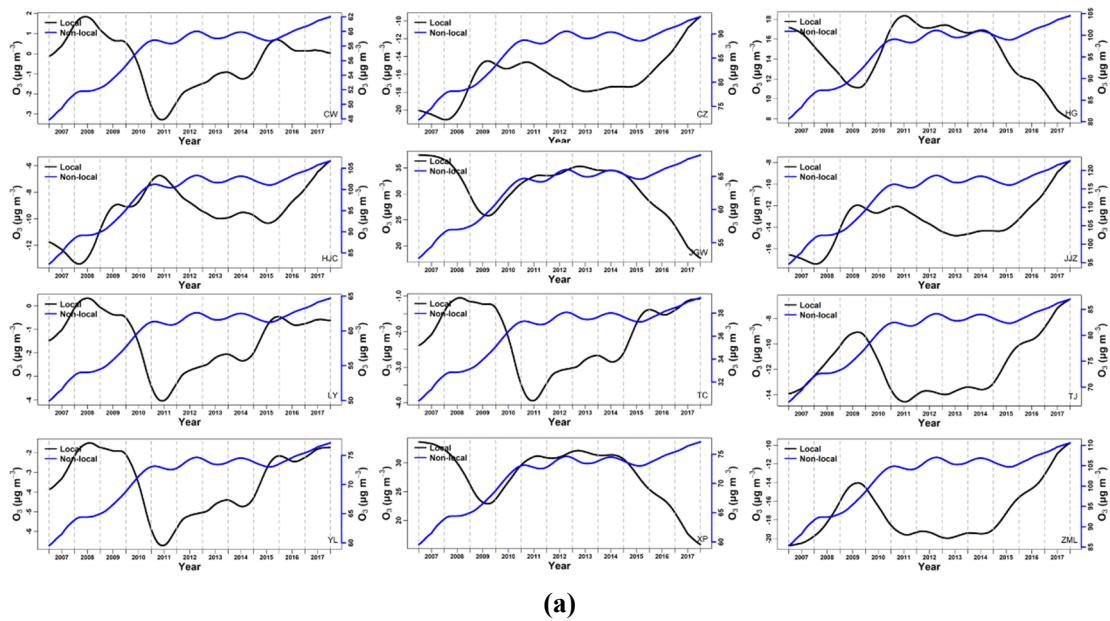
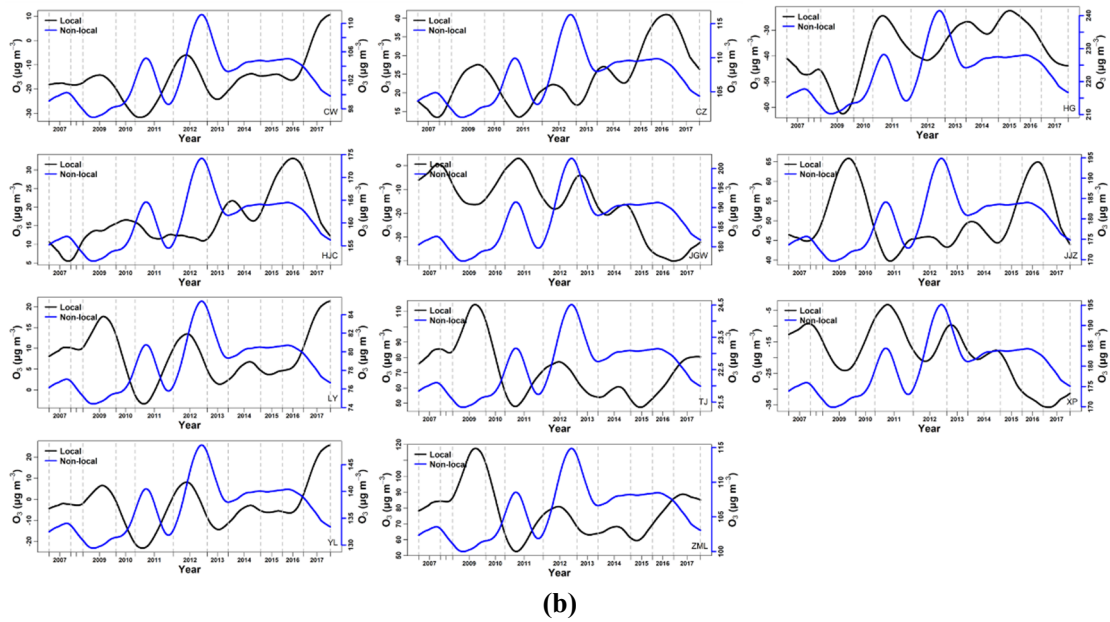


Fig S1. Normalized value of NOx and VOCs emissions over the Pearl River Delta and non-Pearl River Delta area in Guangdong in 2006-2015 against the emissions in 2006 (manuscript under preparation)

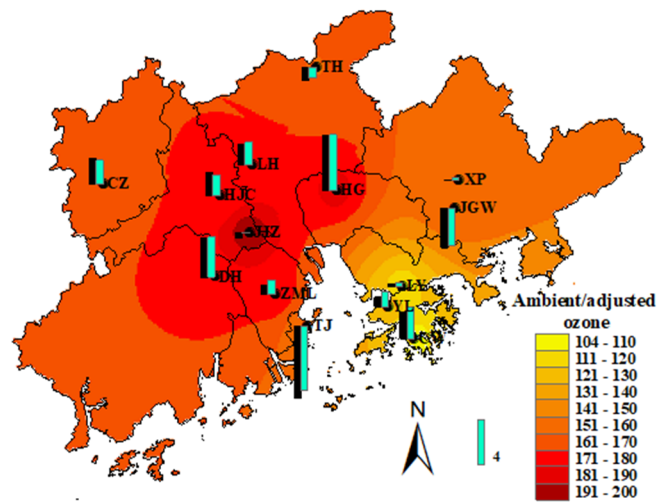
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15 **Fig S2. Long-term trend of ozone contributed by local (black) and non-local (blue) emission sources from 2007 to 2017 during general condition (a) and ozone episodes (b) of different sites.**

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25 **Fig S3. Spatial distribution of averaged ozone concentrations during ozone episodes in the Pearl River Delta and the annual ozone changes ($\mu\text{g m}^{-3} \text{ year}^{-1}$) before and after meteorological adjustment over the fourteen monitoring stations during 2007-2017. The bar length in the legend corresponds to an annual increase of $4 \mu\text{g m}^{-3}$.**

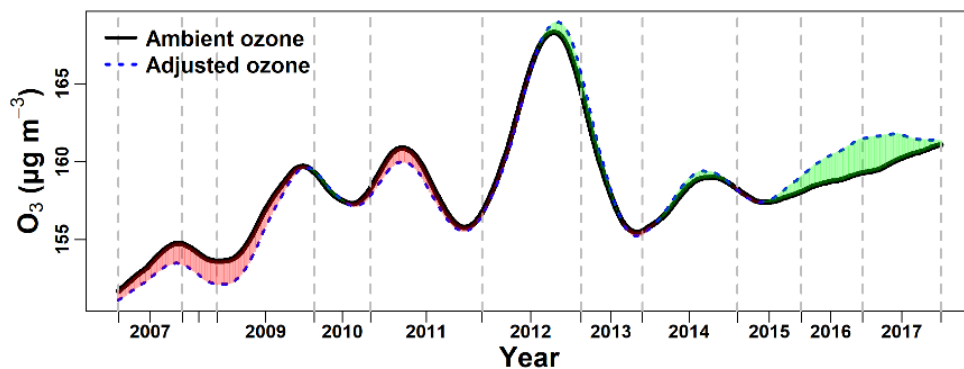


Fig S4. Long-term trends of ambient ozone, meteorologically adjusted ozone, and the meteorological impact in the Pearl River Delta during ozone episodes in 2007-2017.

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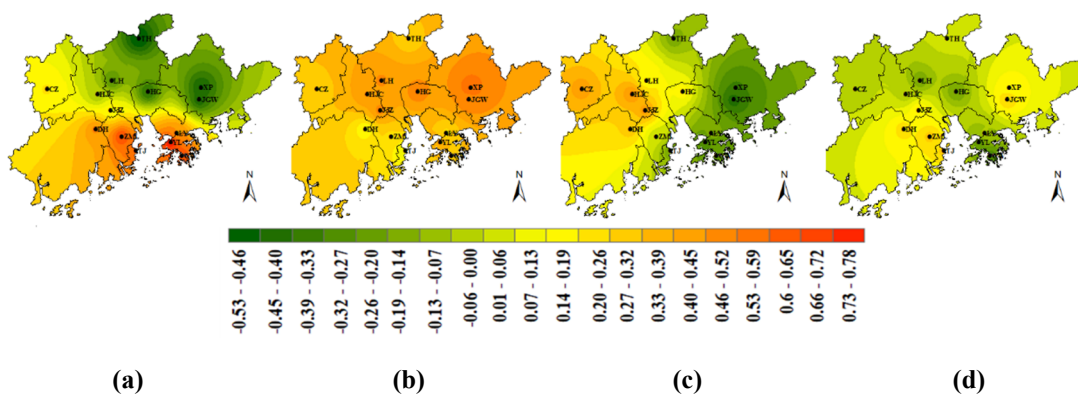


Fig S5. Spatial distribution of principal component loadings (PC1-4: a-d). With positive loadings at all stations, PC2 is assigned to represent impact from non-local emissions. The other three PCs reflect impacts from different local emissions.

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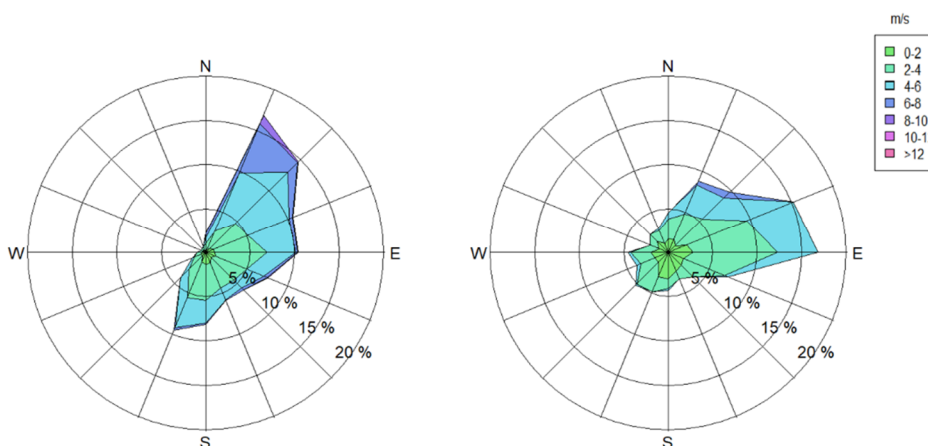


Fig S6. Wind rose under general conditions (left) and during ozone episodes (right) during 2007 and 2017 in the Pearl River Delta.

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