1	Ozone enhancement due to photo-disassociation of
2	nitrous acid in eastern China
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#### 28 Abstract

29 PM<sub>2.5</sub>, a particulate matter with a diameter of 2.5 micrometers or less, is one of the major components of the air pollution in eastern China. In the past few years, China's 30 government made strong efforts to reduce the PM<sub>2.5</sub> pollutions. However, another 31 32 important pollutant (ozone) becomes an important problem in eastern China. Ozone 33  $(O_3)$  is produced by photochemistry, which requires solar radiation for the formation of O<sub>3</sub>. Under heavy PM<sub>2.5</sub> pollution, the solar radiation is often depressed, and the 34 photochemical production of O<sub>3</sub> is prohibited. This study shows that during late 35 spring and early fall in eastern China, under heavy PM<sub>2.5</sub> pollutions, there were often 36 strong O<sub>3</sub> photochemical productions, causing a co-occurrence of high PM<sub>2.5</sub> and O<sub>3</sub> 37 concentrations. This co-occurrence of high PM2.5 and O3 is un-usual and is the main 38 39 focus of this study. Recent measurements show that there were often high HONO surface concentrations in major Chinese mega cities, especially during daytime, with 40 maximum concentrations ranging from 0.5 to 2 ppbv. It is also interesting to note that 41 the high HONO concentrations were occurred during high aerosol concentration 42 43 periods, suggesting that there were additional HONO surface sources in eastern China. Under the high daytime HONO concentrations, HONO can be photo-dissociated to be 44 45 OH radicals, which enhance the photochemical production of O<sub>3</sub>. In order to study the above scientific issues, a radiative transfer model (TUV; Tropospheric 46 Ultraviolet-Visible) is used in this study, and a chemical steady state model is 47 48 established to calculate OH radical concentrations. The calculations show that by 49 including the OH production of the photo-dissociated of HONO, the calculated OH concentrations are significantly higher than the values without including this 50 51 production. For example, by including HONO production, the maximum of OH concentration under the high aerosol condition (AOD=2.5) is similar to the value 52 53 under low aerosol condition (AOD=0.25) in the no-HONO case. This result suggests 54 that even under the high aerosol condition, the chemical oxidizing process for O<sub>3</sub> production can occurred, which explain the co-occurrence of high PM<sub>2.5</sub> and high O<sub>3</sub> 55 in late spring and early fall seasons in eastern China. However, the O<sub>3</sub> concentrations 56 were not significantly affected by the appearance of HONO in winter. This study 57 58 shows that the seasonal variation of solar radiation plays important roles for controlling the OH production in winter. Because the solar radiation is in a very low 59 level in winter, adding the photolysis of HONO has smaller effect in winter than in 60

61	other seasons, and OH remains low values by including the HONO production term.
62	This study provides some important scientific highlights to better understand the $\mathrm{O}_3$
63	pollutions in eastern China.
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65	Keywords; High PM <sub>2.5</sub> and O <sub>3</sub> , eastern China, HONO photolysis
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### 72 **1. Introduction**

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Currently, China is undergoing a rapid economic development, resulting in a higher 74 demand for energy and greater use of fossil fuels. As a result, the high emissions of 75 pollutants produce heavy pollutions in mega cities of eastern China, such as Beijing 76 and Shanghai. For example, in the city of Shanghai (a largest mega city in China), the 77 urban and economical developments of the city are very rapid. During 1990 to 2015, 78 79 the population increased from 13.3 to 24.1 million. The number of automobiles increased from 0.2 million (1993) to 2.0 million (2011). The rapid growing population 80 81 and energy usage caused a rapid increase in the emissions of pollutants, leading to severe air pollution problems in these mega cities (Zhang et al., 2006; Geng et al., 82 83 2007; Deng et al., 2008).

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85 Measurements, such as satellite observations have revealed much higher aerosol pollution in eastern China than in eastern US (Tie et al., 2006). The high aerosol 86 pollution causes a wide range of environmental consequences. Jia et al. (2019) studied 87 Anthropogenic Aerosol Pollution over the Eastern Slope of the Tibetan Plateau, and 88 89 Zhu et al (2018) studied the impact of smoke aerosols from Russian forest fires on the 90 air pollution over Asia. According to a study by Tie et al. (2009a), exposure to extremely high particle concentrations leads to a great increase of lung cancer cases. 91 92 High PM (particular matter) concentrations also significantly reduce the range of 93 visibility in China's mega cities (Deng et al., 2008). According to a recent study, the high aerosol pollution causes important effects on the crop (rice and wheat) 94 95 production in eastern China (Tie et al., 2016).

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97 In the troposphere, ozone formation is resulted from a complicated chemical process, and requires ozone precursors, such as VOCs (volatile organic carbons) and  $NO_X$  = 98 NO + NO<sub>2</sub> (nitrogen oxides) (Sillman, 1995). As the increase in industrial activity and 99 number of automobiles, the precursors of ozone (O<sub>3</sub>) and the global budget of 100 oxidization are also significantly increased (Huang et al., 2017; Huang et al., 2018). 101 As a result, O<sub>3</sub> pollution becomes a serous pollution problem in Shanghai and other 102 Chinese mega cities (Geng et al., 2010; Tie 2009b; Tie et al., 2015). The effects on O<sub>3</sub> 103 production rate can be characterized as either NO<sub>X</sub>-sensitive or VOC-sensitive 104

105 conditions. For the city areas,  $O_3$  production is generally VOC-sensitive, while in the 106 remote area,  $O_3$  production is generally NOx-sensitive in eastern China (Sillman, 107 1995; Zhang et al., 2003; Lei et al., 2004; Tie et al., 2013). Thus, better understanding 108 the trends of  $O_3$  precursors (VOCs, NO<sub>X</sub>) is important to determine the  $O_3$  trends in 109 Shanghai (as well as many large cities in China).

In the past few years, China's government made strong efforts to reduce the PM<sub>2.5</sub> 110 pollutions. However, another important pollutant (O<sub>3</sub>) becomes an important problem 111 in eastern China. Several studies regarding the O<sub>3</sub> formation are previously studied in 112 113 Shanghai. For example, Geng et al. (2007; 2008) study the relationship between O<sub>3</sub> precursors (NOx and VOCs) for the ozone formation in Shanghai. Tie et al. (2009) 114 115 study the short-term variability of O<sub>3</sub> in Shanghai. Their study suggested that in addition to the ozone precursors, meteorological conditions, such as regional transport, 116 have also strong impacts on the ozone concentrations. During September 2009, a 117 major field experiment (the MIRAGE-Shanghai) was conducted in Shanghai, and 118 119 multiply chemical species were measured during the experiment. The summary of the measurements by Tie et al (2013) suggests that the ozone formation in Shanghai is 120 121 under VOC-sensitive condition. However, if the emission ration of NOx/VOCs reduces to a lower value (0.1-0.2), the ozone formation in Shanghai will switch from 122 123 VOC-sensitive condition to NOx-sensitive condition.

Despite of some progresses have been made for the ozone formation in mega cities in 124 China, it is still lack of study of ozone development in large cities of China. For 125 126 example, this study shows that during late spring and early fall in eastern China, under heavy PM<sub>2.5</sub> pollutions, there were often strong O<sub>3</sub> chemical productions, 127 128 causing the co-occurrence of high PM<sub>2.5</sub> and O<sub>3</sub> concentrations. Under heavy aerosol 129 condition, the solar radiation is depressed, significantly reducing the photochemical 130 production of O<sub>3</sub>. This co-occurrence of high PM<sub>2.5</sub> and O<sub>3</sub> is an unusual and is the focus of this study. He and Carmichael (1999) suggest that aerosol particles can 131 132 enhance the scattering of solar radiation, enhancing the flux density inside the boundary layer. Recent measurements also show that there were often high HONO 133 134 concentrations in major Chinese mega cities, especially during daytime, with maximum concentrations ranging from 0.5 to 2 ppbv (Huang et al., 2017). Zhang et al. 135 (2016) suggest that there are several potential HONO sources, including surface 136

emissions, conversion of NO<sub>2</sub> at the ocean surface, etc., and adding these sources can
improve the calculated HONO concentrations. It is also interesting to note that the
high HONO surface concentrations were occurred during high aerosol concentration
periods, suggesting that there are additional HONO surface sources in eastern China.
Under the high daytime HONO concentrations, HONO can be photo-dissociated to be
OH radicals, which enhance the photochemical production of O<sub>3</sub>.

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The paper is organized as follows: in Section 2, we describe the measurement of  $O_3$ and  $PM_{2.5}$ . In Section 3, we describe the calculation of photo-dissociated rate of HONO and a steady state model for the calculation of OH, and the causes of high  $O_3$ production under the heavy aerosol condition. Section 4 shows a brief conclusion of the results.

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## 2. Measurements of O<sub>3</sub> and PM<sub>2.5</sub>

There are long-term measurements in Eastern China by Chinese Environment 152 Protection Agency (CEPA) for monitoring the air quality in China. In eastern China, 153 especially in the capital city of China (Beijing), there are often heavy air pollutions, 154 especially for fine particular matter (PM<sub>2.5</sub> – the radium of particle being less than 2.5 155 um). Figure 1 shows the measurement sites in Beijing, in which the measured 156 concentrations of PM<sub>2.5</sub> and O<sub>3</sub> are used to the analysis. In the region, the air 157 pollutions were very heavy, especially in winter (Long et al., 2016; Tie et al., 2017). 158 The previous studies suggested that the both aerosol and  $O_3$  pollutions became the 159 160 major pollutants in the region (Li et al., 2017).

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162 Figure 2 shows the daily averaged concentrations of PM<sub>2.5</sub> and O<sub>3</sub> in the Beijing region in 2015. The daily averaged concentrations show that there were strong daily 163 and seasonal variations for both the concentrations of PM2.5 and O3. Despite the daily 164 variation, the concentrations of PM<sub>2.5</sub> existed a strong seasonal variation. For example, 165 there were very high concentrations during winter, with maximum of  $\sim 300 \ \mu g/m^3$ . 166 While in summer, the maximum concentrations reduced to  $\sim 150 \ \mu g/m^3$ . The seasonal 167 168 variability of O<sub>3</sub> concentrations were opposite with the PM<sub>2.5</sub> concentrations, with lower concentrations in winter (< 50  $\mu$ g /m<sup>3</sup>) and higher concentrations in summer (> 169

150  $\mu$ g/m<sup>3</sup>). These seasonal variations of PM<sub>2.5</sub> and O<sub>3</sub> have been studied by previous 170 studies (Tie and Cao, 2017; Li et al., 2017). Their results suggest that the winter high 171 PM<sub>2.5</sub> concentrations were resulted from the combination of both the high emissions 172 (heating season in the Beijing region), and poor meteorological ventilation conditions, 173 174 such as lower PBL (Planetary Boundary Layer) height (Quan et al., 2013; Tie et al. 2015). According to the photochemical theory of  $O_3$  formation, the summer high and 175 winter low O<sub>3</sub> concentrations are mainly due to seasonal variation of the solar 176 radiation (Seinfeld, J. H. and Pandis, 2006). 177

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179 The heavy aerosol concentrations play important roles to reduce solar radiation, causing the reduction of O<sub>3</sub> formation. (Bian et al., 2007). As we show in Fig. 3 180 (upper panel), during wintertime, the O<sub>3</sub> concentrations were strong anti-correlated 181 with the PM<sub>2.5</sub> concentrations, suggesting that the reduction of solar radiation by 182 aerosol particles have important impact on the reduction of O<sub>3</sub> concentrations. Figure 183 3 (upper panel) also shows that the relationship between O<sub>3</sub> and PM<sub>2.5</sub> was not 184 linearly related. For example, when the concentrations of PM25 were less than 100 185  $\mu$ g/m<sup>3</sup>, O<sub>3</sub> concentrations rapidly decreased with the increase of PM<sub>2.5</sub> concentrations. 186 187 In contrast, when the concentrations of PM<sub>2.5</sub> were greater than 100  $\mu$ g/m<sup>3</sup>, O<sub>3</sub> concentrations slowly decreased with the increase of PM<sub>2.5</sub> concentrations. This is 188 consistent with the result of Bian et al (2007). 189

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191 It is interesting to note that from late spring to early fall periods, the correlation between PM<sub>2.5</sub> and O<sub>3</sub> concentrations was positive relationship compared to the 192 negative relationship in winter (see Fig. 3 (lower panel)). This result suggests that O<sub>3</sub> 193 production was high during the heavy haze period, despite the solar radiation was 194 greatly depressed. In order to clearly display this unusual event, we illustrate diurnal 195 variations of PM<sub>2.5</sub> and O<sub>3</sub> and NO<sub>2</sub> during a fall period (from Oct.5 to Oc. 6, 2015). 196 Figure 4 shows that during this period (as a case study), the PM<sub>2.5</sub> concentrations were 197 very high, ranging from 150 to 320  $\mu$ g/m<sup>3</sup>. Under such high aerosol condition, the 198 199 solar radiation should be significantly reduced, and O<sub>3</sub> photochemical production would be reduced. However, the diurnal variation of O<sub>3</sub> was unexpectedly strong, 200 with high noontime concentration of >220  $\mu$ g/m<sup>3</sup> and very low nighttime 201 concentration of ~25  $\mu$ g/m<sup>3</sup>. This strong diurnal variation was due to the 202 photochemical activity, which suggested that during relatively low solar conditions, 203

the photochemical activities of  $O_3$  production was high. According to the theory of the  $O_3$  chemical production, the high  $O_3$  production is related to high oxidant of OH (Seinfeld and Pandis, 2006), which should not be occurred during lower solar radiation. This result brings important issue for air pollution control strategy, because both PM<sub>2.5</sub> and O<sub>3</sub> are severe air pollutants in eastern China.

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To clearly understand the effect of the high aerosol concentrations on solar radiation, 210 we investigate the meteorological conditions, such as cloud covers, relation humidity 211 212 (RH), and solar radiation during the period of the case study (see Figs. 5 and 6). Figure 5 shows that the cloud condition was close to the cloud free condition, but 213 214 there was a very heavy aerosol layer in the Beijing region, suggesting that cloud cover played a minor role in the reduction of the solar radiation. The measured RH values 215 (not shown) were generally higher than 60%, with a maximum of 95% during the 216 period. As a result, the high aerosol concentrations companied by high RH produced 217 important effects on solar radiation. As shown in Fig. 6, the daytime averaged solar 218 radiation was significantly reduced (about 40% reduction in Oct. 5-6 period compared 219 with the value of Oct. 8). 220

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## **3. Method**

In order to better understand the  $O_3$  chemical production occurred in heavy aerosol condition in eastern China, the possible  $O_3$  production in such condition is discussed. Ozone photochemical production (P[O<sub>3</sub>]) is strongly related to the amount of OH radicals (Chameides et al., 1999). According to the traditional theory, the amount of surface OH radicals is proportional to the surface solar radiation, which is represented by

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$$[OH] = P[HOx]/L[HOx]^*$$
 (R-1)

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Where [OH] represents the concentration of hydroxyl radicals ( $\#/cm^3$ ); HOx represents the concentration of HO<sub>2</sub> + OH ( $\#/cm^3$ ); P[HOx] represents the photochemical production of HOx ( $\#/cm^3/s$ ); and L[HOx]\* (1/s) represents the photochemical destruction of HOx, which is normalized by the concentrations of OH. The major process for the photochemical production of P[HOx] is through the O<sub>3</sub> photolysis and follows by the reaction with atmospheric water vapor. It can be expressed as

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$$P[HOx] = J_1[O_3]/(k_1 \times am) \times 2.0 \times k_2[H_2O] = P_1[HOx]$$
(R-2)

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Where J<sub>1</sub> represents the photolysis of  $O_3 + hv \rightarrow O^1D$ ; k<sub>1</sub> represents the reaction rate 243 of  $O^1D + am \rightarrow O^3P$ ; and k<sub>2</sub> represents the reaction rate of  $O^1D + H_2O \rightarrow 2OH$ . As 244 we can see, this HOx production is proportional to the magnitude of solar radiation 245 246  $(J_1)$ , and  $J_1$  is the O<sub>3</sub> photolysis with the solar radiation. Figure 7 shows the relationship between the values of J<sub>1</sub> and aerosol concentrations in October at 247 middle-latitude calculated by the TUV model (Madronich and Flocke, 1999). This 248 result suggests that under the high aerosol concentrations (AOD = 2.5), the J<sub>1</sub> value is 249 strongly depressed, resulting in significant reduction of OH concentrations and O<sub>3</sub> 250 production. For example, the maximum  $J_1$  value is about 2.7x10<sup>-5</sup> (1/s) with lower 251 aerosol values (AOD = 0.25). According to the previous study, the surface PM<sub>2.5</sub> 252 concentrations were generally smaller than 50  $\mu$ g/m<sup>3</sup> with this AOD value (Tie et al., 253 2017). However, when the AOD value increase to 2.5 (the  $PM_{2.5}$  concentrations are 254 generally >100  $\mu$ g/m<sup>3</sup>), the maximum J<sub>1</sub> value rapidly decreases to about 6x10<sup>-6</sup> (1/s), 255 which is about 450% reduction compared to the value with AOD=0.25. This study 256 suggests that under high PM<sub>2.5</sub> concentrations (>100  $\mu$ g/m<sup>3</sup>), the photochemical 257 production of OH (P[HOx]) is rapidly decreased, leading to low OH concentrations, 258 259 which cannot initiate the high oxidation of O<sub>3</sub> production. As a result, the high O<sub>3</sub> production shown in Fig. 4 cannot be explained. Other sources for O<sub>3</sub> oxidation are 260 needed to explain this result. 261

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Recent studies show that the HONO concentrations are high in eastern China (Huang 263 et al., 2017). Because under high solar radiation, the photolysis rate of HONO is very 264 high, resulting in very low HONO concentrations in daytime (Seinfeld and Pandis, 265 2006). These measured high HONO concentrations are explained by their studies. 266 One of the explanations is that there are high surface HONO sources during daytime, 267 which produces high HONO concentrations (Huang et al., 2017). Zhang et al. (2016) 268 suggest that there are several potential HONO sources, including surface emissions, 269 conversion of NO<sub>2</sub> at the ocean surface, etc. Zhang et al. (2016) parameterized these 270

potential HONO sources in the WRF-Chem model, and the calculated HONOconcentrations are increased in the WRF-Chem model.

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The version of the WRF-Chem model is based on the version developed by Grell et al. 274 275 (2015), and is improved mainly by Tie et al. (2017) and Li et al. (2011). The chemical mechanism chosen in this version of WRF-Chem is the RADM2 (Regional Acid 276 Deposition Model, version 2) gas-phase chemical mechanism. For the calculation of 277 HONO, only the gas-phase chemistry of OH+NO is included to calculate HONO 278 279 concentrations. As shown in Fig. 8, the calculated HONO concentrations are 280 significantly smaller than the measured HONO values in eastern China, suggesting that in addition to the gas-reaction, there are missing HONO sources (surface sources 281 282 or others). Because these missing sources are not fully understood and large uncertainty is remained, in the following calculation, we compare the OH 283 284 concentrations due to both calculated HONO (without the missing sources) and the measured HONO concentrations to illustrate the importance of these missing sources 285 286 for the production of OH radicals and to suggest that further study to better understand the missing sources is an urgent scientific issue. 287

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289 Figure 8 shows the measured HONO concentrations in three large cities in China 290 (Shanghai, Xi'an, and Beijing) during fall and winter. It also shows the corresponding PM<sub>2.5</sub> and O<sub>3</sub> in the 3 cities (i.e., Fig. 8a for Beijing, Fig. 8b for Shanghai, and Fig. 8c 291 292 for Xian). It shows that the measured HONO concentrations were high, ranging from 293 sub-ppby to a few ppby, with higher values during morning, and lower values in daytime. The co-occurrences of high PM<sub>2.5</sub> and O<sub>3</sub> happened in the 3 cities. As a 294 result, we think that the high HONO is a common event in large cities in eastern 295 China, especially in daytime. This high HONO is also measured by previous studies 296 (Zhang et al. 2016; Huang et al. 2017). In this study, we make an assumption that the 297 298 co-occurrence between O<sub>3</sub> and PM<sub>2.5</sub> occurred under high HONO concentrations. We 299 note that using this assumption may result in some uncertainties in estimating the 300 effect of HONO on OH. For example, using the measured HONO in Xi'an and Beijing could produce 1-2 times higher OH production by photolysis of HONO than 301 the result by using the data from Shanghai. In this case, we use the measured HONO 302 from Shanghai to avoid the over estimate of the HONO effect, which can be 303 considered as a low-limit estimation. 304

It is also interesting to note that the high HONO concentrations were occurred during high aerosol concentration periods. Figure 9 illustrates that when the PM<sub>2.5</sub> concentrations increased to 70-80  $\mu$ g/m<sup>3</sup>, and the HONO concentrations enhanced to 1.4-18 ppbv during September in Shanghai. This measured high HONO concentrations were significantly higher than the calculated concentrations (shown in Fig. 8), suggesting that some additional sources of HONO are needed. This result is consistent with the HONO measurements in other Chinese cities (Huang et al. 2017).

The high HONO concentrations in daytime become a significant source of OH radicals. As a result, the OH production rate (P[HOx]) can be written to the following reactions.

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Because the chemical lifetime of OH is less than second, OH concentrations can be calculated according to equilibrium of chemical production and chemical loss. With the both OH chemical production processes, the OH concentrations can be calculated by the following equation (Seinfeld and Pandis, 2006).

 $= J_1[O_3]/(k_1 \times am) \times 2.0 \times k_2[H_2O] + J_2 \times [HONO]$  (R-4)

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327 P1 + P2 =

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$$P1 + P2 = L1 + L2$$

 $P_2[HOx] = J_2 \times [HONO]$ 

 $P[HOx] = P_1[HOx] + P_2[HOx]$ 

Where P1 and P2 are the major chemical productions, expressed in R-4, and L1 andL2 are the major chemical loss of OH, and represent by

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333 L2: 
$$HO_2 + HO_2 \rightarrow H_2O_2 + O_2$$
 (R-6)

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Under high NOx condition, such as in the large cities in eastern China, NOx
concentrations were often higher to 50 ppbv (as shown in Fig. 4). As a result, the L1
term is larger than L2. The OH concentrations can be approximately expressed as

(R-3)

(R-5)

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$$[HO] = \{J_1[O_3]/(k_1 \times am) \times 2.0 \times k_2[H_2O] + J_2 \times [HONO]\}/$$

k<sub>3</sub>[NO<sub>2</sub>]

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Where  $k_3$  is the reaction coefficient of OH + NO<sub>2</sub>  $\rightarrow$  HNO<sub>3</sub>.

344 **4. Result and analysis** 

## 346 **<u>4.1. OH productions in different HONO conditions</u>**

- In order to quantify the individual effects of these two OH production terms (P1 and 348 P2) on the OH concentrations, the P1 and P2 are calculated under different daytime 349 350 HONO conditions (calculated low HONO and measured high HONO concentrations). Figure 10 shows that under the low HONO condition, the P1 is significantly higher 351 than P2, and P2 has only minor contribution to the OH values. For example, the 352 maximum of P1 occurred at 13 pm, with a value of  $65 \times 10^6$  #/cm<sup>3</sup>/s. In contrast, the 353 maximum of P2 occurred at 10 am, with a value of  $15 \times 10^6$  #/cm<sup>3</sup>/s. However, under 354 high HONO condition, the P2 plays very important roles for the OH production. The 355 maximum of P2 occurred at 11 am, with a value of  $350 \times 10^6$  #/cm<sup>3</sup>/s, which is about 356 500% higher than the P1 value. It is important to note that this calculation is based on 357 the high aerosol condition (AOD = 2.5) in September. This result can explain the high 358 359 O<sub>3</sub> chemical production in Fig. 4.
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#### 4.2. OH in different aerosol conditions

In order to understand the effect of aerosol conditions, especially high aerosol 363 conditions, on the OH concentrations. Figure 11 shows the OH concentrations with 364 and without HONO production of OH. With including the HONO production (i.e., 365 including P1 and P2), the calculated OH concentrations are significantly higher than 366 without including this production (i.e., only including P1). The both calculated OH 367 concentrations are rapidly changed with different levels of aerosol conditions. For 368 example, without HONO production, the maximum OH concentration is about 369  $7.5 \times 10^5$  #/cm<sup>3</sup> under low aerosol condition (AOD=0.25). In contrast, the maximum 370 OH concentration rapidly reduced to  $1.5 \times 10^5$  #/cm<sup>3</sup> under high aerosol condition 371 (AOD=2.5), and further decreased to  $1.0 \times 10^5 \, \text{#/cm}^3$  with the AOD value of 3.5. In 372 contrast, with including HONO production, the OH concentrations significantly 373

(R-5)

increased. Under higher aerosol condition (AOD=2.5), the maximum of OH concentration is about  $7.5 \times 10^5$  #/cm<sup>3</sup>, which is the same value under low aerosol condition in the no-HONO case. This result suggests that the measured high O<sub>3</sub> production occurred in the high aerosol condition is likely due to the high HONO concentrations in Shanghai.

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## 380 **<u>4.3. Effects of clouds</u>**

Cloud cover can have very important impacts on the photolysis of HONO, which can affect the effect of HONO on the OH radicals. The above calculations are based on the cloud-free condition, with heavy aerosol concentration in the Beijing region. As shown in Fig. 5, during the case study period (Oct 5 to 6, 2015) (see Fig. 4), the weather map shows that the cloud-free condition, with heavy aerosol condition.

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In order to understand the effects of cloud on the photolysis of HONO, we include 388 389 different cloud covers in the TUV model. The calculated results show in Fig. 12. The results show that the thin cloud (with cloud cover in 2 km and cloud water of 10 390  $g/m^3$ ), could reduce the photolysis rate of HONO by about 40%, but the HONO could 391 still remain important effects. However, with dense cloud condition (with cloud 392 covers at 2 and 3 km and cloud water of 50 10  $g/m^3$ ), the photolysis rate of HONO 393 could reduce by 9-10 times by the cloud. In this case, adding photolysis rate of 394 HONO cannot produce important effect on OH radicals and the production of O<sub>3</sub>. 395

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## 397 <u>4.4. OH in winter</u>

399 The measurement of  $O_3$  also shows that the concentrations in winter were always low (see Fig. 2), suggesting that the  $O_3$  concentrations were not significantly affected by 400 the appearance of HONO. Figure 13 shows the OH concentrations in September and 401 December. It shows that under different aerosol conditions, OH concentrations in 402 December were very low compared with the values in September. Both the calculated 403 OH concentrations include the HONO production term. For example, under the 404 condition of AOD=2.5, the maximum OH is about  $7.5 \times 10^5$  #/cm<sup>3</sup> in September, while 405 it rapidly reduces to  $1.5 \times 10^5$  #/cm<sup>3</sup> in December. Under the condition of AOD=3.5, 406 the maximum OH is still maintaining to a relative high level  $(4.5 \times 10^5 \text{ } \#/\text{cm}^3)$  in 407 September. However, the maximum OH values are extremely low in December, with 408

409 maximum value of  $0.5 \times 10^5$  #/cm<sup>3</sup> in December. Because both the OH chemical 410 productions (P1 and P2) are strongly dependent upon solar radiation (see equation 411 R-4), the seasonal variation of solar radiation plays important roles for controlling the 412 OH production in winter (see Fig. 13). Because the solar radiation is in a very low 413 level in winter, adding the photolysis of HONO has smaller effect in winter than in 414 other seasons and OH remains low values by including the HONO production term.

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#### 416 Summary

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Currently, China is undergoing a rapid economic development, resulting in a high 418 419 demand for energy, greater use of fossil fuels. As a result, the high emissions of pollutants produce heavy aerosol pollutions (PM<sub>2.5</sub>) in eastern China, such as in the 420 421 mega city of Beijing. The long-term measurements show that in addition to the heavy aerosol pollution, the O<sub>3</sub> pollution becomes another major pollutants in the Beijing 422 423 region. The measured results show that there were very strong seasonal variation in the concentrations of both PM<sub>2.5</sub> and O<sub>3</sub> in the region. During winter, the seasonal 424 425 variability of O<sub>3</sub> concentrations were anti-correlated with the PM<sub>2.5</sub> concentrations. However, from late spring to early fal, the correlation between PM<sub>2.5</sub> and O<sub>3</sub> 426 427 concentrations was positive compared to the negative in winter. This result suggests 428 that during heavy aerosol condition (the solar radiation was depressed), the O<sub>3</sub> chemical production was still high, appearing a co-occurrence of high PM2.5 and O3 in 429 some cases from late spring to early fall. This co-occurrence of high PM<sub>2.5</sub> and O<sub>3</sub> is 430 the focus of this study. The results are highlighted as follows; 431

432

(1) There are high daytime HONO concentrations in major Chinese mega cities, such
as in Beijing and Shanghai. It is also interesting to note that the high HONO
concentrations were occurred during high aerosol concentration periods. Under
the high daytime HONO concentrations, HONO can be photo-dissociated to be
OH radicals, and becomes an important process to produce OH.

438 (2) With including the OH production of measured HONO concentrations, the 439 calculated OH concentrations are significantly higher than without including this 440 production. For example, without HONO production, the maximum OH 441 concentration is about  $7.5 \times 10^5$  #/cm<sup>3</sup> under low aerosol condition (AOD=0.25), 442 and rapidly reduced to  $1.5 \times 10^5$  #/cm<sup>3</sup> under high aerosol condition (AOD=2.5) in 443 September. In contrast, by including HONO production, the OH concentrations 444 significantly increased. For example, under higher aerosol condition (AOD=2.5), 445 the maximum of OH concentration is about  $7.5 \times 10^5$  #/cm<sup>3</sup>, which is similar to the 446 value under low aerosol condition in the no-HONO case. This result suggests that 447 even under the high aerosol conditions, the chemical oxidizing process for O<sub>3</sub> 448 production can be active. This result is likely for explaining the co-occurrence of 449 high PM<sub>2.5</sub> and high O<sub>3</sub> from late spring to early in eastern China.

(3) The measurement of O<sub>3</sub> also shows that the concentrations in winter were always
low, suggesting that the O<sub>3</sub> concentrations were not significantly affected by the
appearance of HONO. The calculated result shows that the seasonal variation of
solar radiation plays important roles for controlling the OH production in winter.
Because the solar radiation is in a very low level in winter, adding the photolysis
of HONO has smaller effect in winter than in other seasons, and OH remains low
values by including the HONO production term.

In recent years, the  $PM_{2.5}$  pollutions are reduced due to the large control efforts by the Chinese government, the  $O_3$  pollutions become another severe pollution problem in eastern China. This study is important, because it provides some important scientific highlights to better understand the  $O_3$  pollutions in eastern China.

461

462 Data availability. The data used in this paper can be provided upon request from
463 Xuexi Tie (tiexx@ieecas.cn).

464

465 **Author contributions**. XT came up with the original idea of investigating the 466 scientific issue. XT and JX designed the analysis method. XL, GL and SZ provided 467 the observational data and helped in discussion. XT prepared the manuscript with 468 contributions from all co-authors.

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# 592 **Figure Caption**

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**Fig. 1.** The geographic locations of the measurement sites in Beijing, in which the measured concentrations of  $PM_{2.5}$  and  $O_3$  are used to the analysis.

**Fig. 2.** The daily averaged concentrations of  $PM_{2.5}$  and  $O_3$  in the Beijing region in 2015. The concentrations are averaged over all sites shown in Fig. 1. The blue lines represent the  $PM_{2.5}$  concentrations ( $\mu g/m^3$ ), and the red bars represent the  $O_3$ concentrations ( $\mu g/m^3$ ). The rectangles show some typical events during winter (green), spring and fall (orange), and summer (red).

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**Fig. 3.** The correlation between  $O_3$  and  $PM_{2.5}$  concentrations during winter (upper panel) and from late spring to eraly fall (lower panel). During winter,  $O_3$ concentrations were strong anti-correlated with the  $PM_{2.5}$  concentrations. From late spring to early fall,  $O_3$  concentrations were correlated with the  $PM_{2.5}$  concentrations.

Fig. 4. The diurnal variations of PM<sub>2.5</sub> (blue line) and O<sub>3</sub> (red line), and NO<sub>2</sub> (green line) during a fall period (from Oct.5 to Oc. 6, 2015). It shows that with high PM<sub>2.5</sub> condition, there was a strong O<sub>3</sub> diurnal variation.

Fig. 5. The cloud condition during the period of the case study (between Oct 5 and 6, 2015) in the Beijing region. The bright white color shows the cloud covers, and the grey white shows the haze covers. The Beijing region was under the heavy haze conditions during the period.

Fig. 6. The measured solar radiation  $(W/m^2)$  from Oct. 3 to Oct. 9, 2015 in Beijing. The upper panel shows hourly values, and the lower panel shows the daytime averaged values.

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Fig. 7. The effect of aerosol levels with AOD = 0.25 (black line), AOD = 2.5 (red line), AOD = 3.5 (blue line), and AOD = 4.0 (green line) on the  $O_3$  photolysis calculated by the TUV model in October at middle-latitude.

Fig. 8a. The measured HONO concentrations (ppbv) and the  $PM_{2.5}$  and  $O_3$  daily concentrations in Beijing. The upper panel shows the measured daily concentrations of  $PM_{2.5}$  and  $O_3$  as shown in Fig.2. The dark-red line was measured HONO in Beijing from 1 to 27 January, 2014.

630 **Fig. 8b.** The measured HONO concentrations (ppbv) and the  $PM_{2.5}$  and  $O_3$  daily 631 concentrations in Shanghai. The upper panel shows the measured daily concentrations 632 of  $PM_{2.5}$  and  $O_3$  in 2015. The dark-red line was measured in Shanghai from 9 to 18 633 September, 2009. The green line was calculated by the WRF-Chem model.

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Fig. 8c. The measured HONO concentrations (ppbv) and the  $PM_{2.5}$  and  $O_3$  daily concentrations in Xi'an. The upper panel shows the measured daily concentrations of  $PM_{2.5}$  and  $O_3$  in 2015. The red line was measured HONO in Xi'An from 24 July to August 6, 2015.

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640 **Fig. 9.** The measured HONO (upper left panel),  $PM_{2.5}$  concentrations (lower left 641 panel), and O<sub>3</sub> concentrations (upper right panel) in fall in Shanghai. It illustrates that

- the high HONO concentrations were corresponded with high  $PM_{2.5}$  concentrations.
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**Fig. 10.** The calculated OH production P(HOx) (#/cm<sup>3</sup>/s) by using the model calculated HONO (low concentrations) (in the upper panel) and by using the measured HONO (high concentrations) (in the lower panel). The red bars represent the calculation of the P1 term, and the red bars represent the calculation of the P2 term (OH production from HONO).

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**Fig. 11.** The calculated OH concentrations (#/cm<sup>3</sup>) with (upper panel) and without (lower panel) HONO production of OH, under different aerosol levels. Dark red (AOD=0.25), red (AOD=2.5)), red (AOD=3.5)), and red (AOD=4.0).

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**Fig. 12.** The effect of cloud cover on the photolysis rate of HONO (J[HONO]). The blue, red, and green lines represent the cloud water vapor of 0 (cloud-free),  $10 (g/m^3$ thin cloud), and 50 (g/m<sup>3</sup> – thick cloud), respectively. The left panel (A) represents the light aerosol condition, with AOD of 0.25, and the right panel (B) represents the heavy aerosol condition, with AOD of 2.5.

Fig. 13. The calculated OH concentrations in September (blue bars) and December(dark red bars), under different aerosol levels.

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# Figures

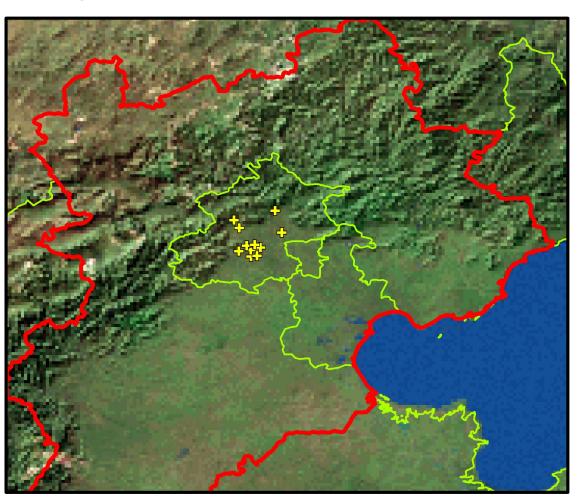
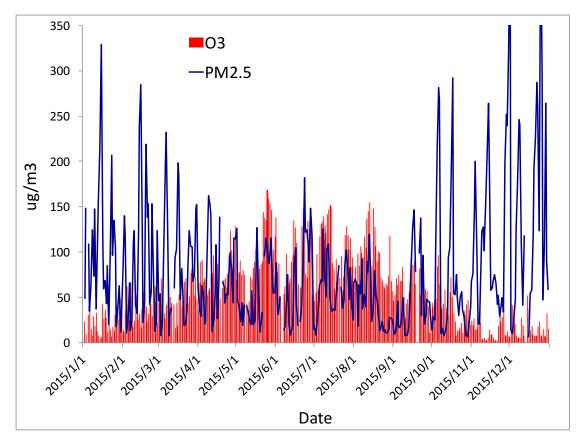
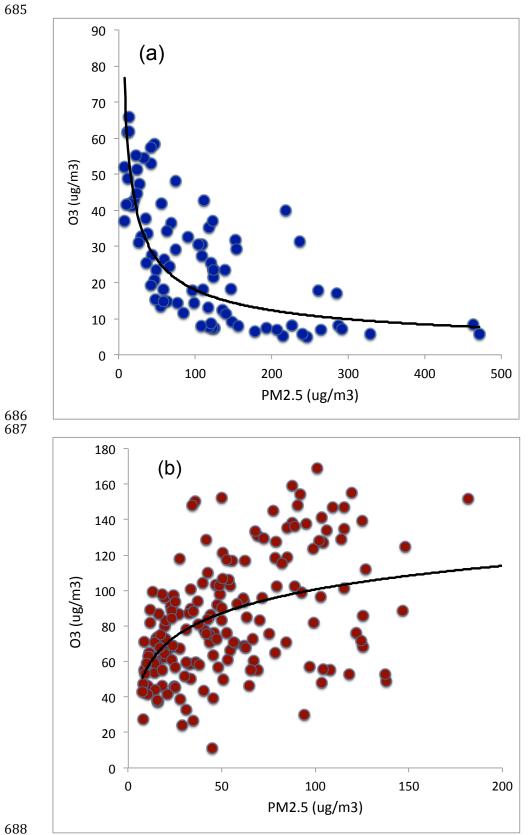


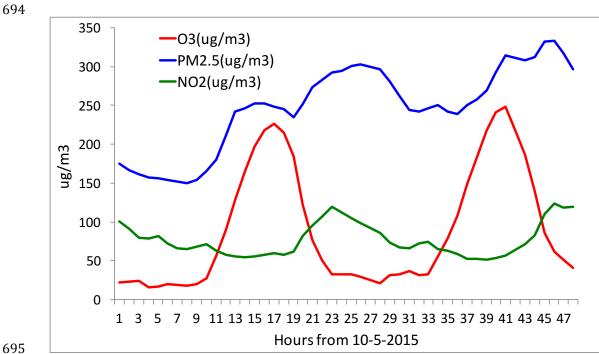
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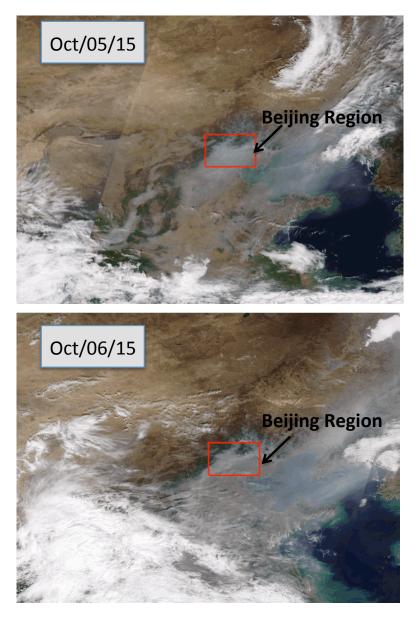
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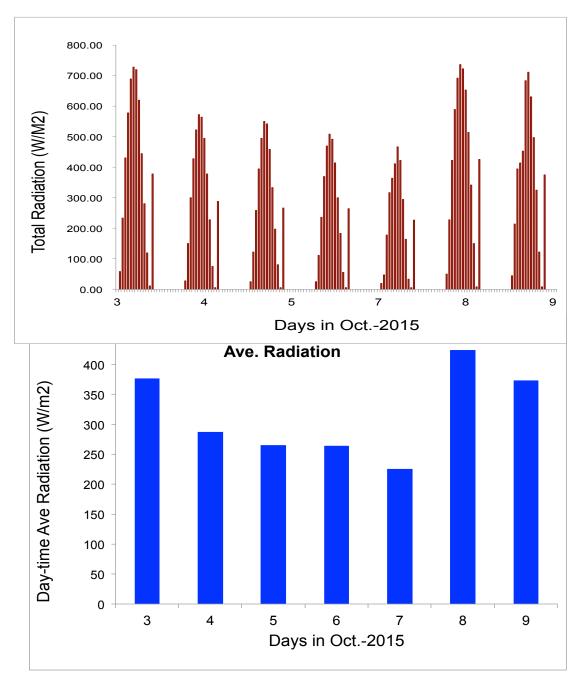
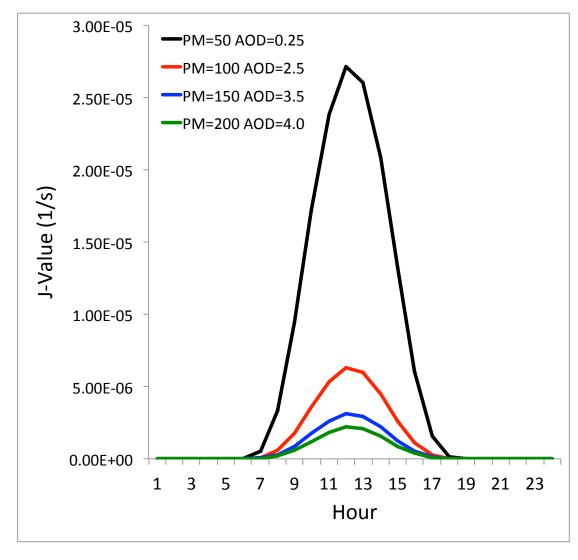
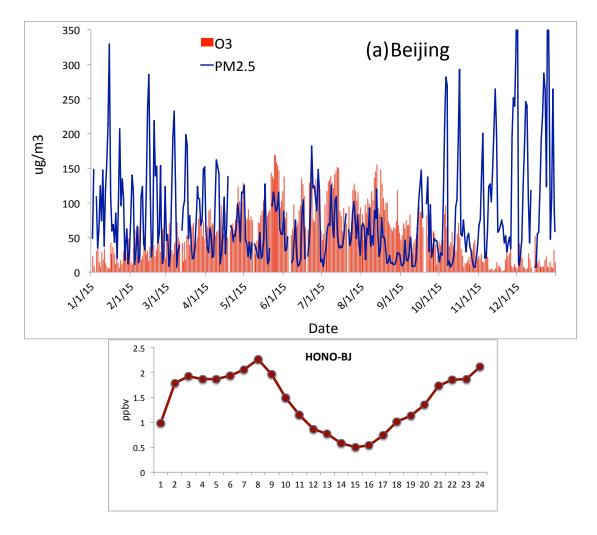




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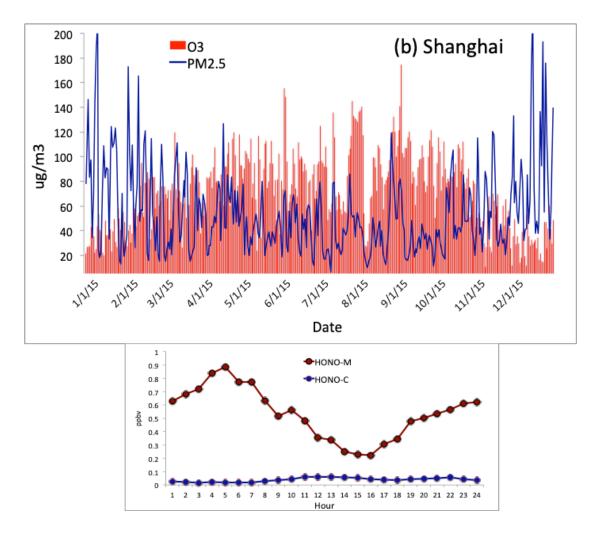
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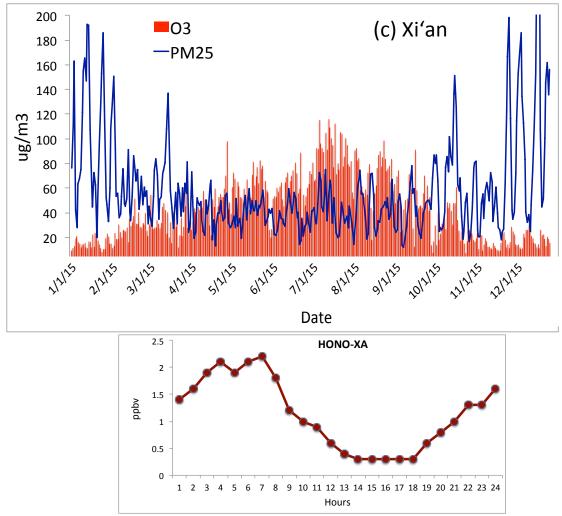


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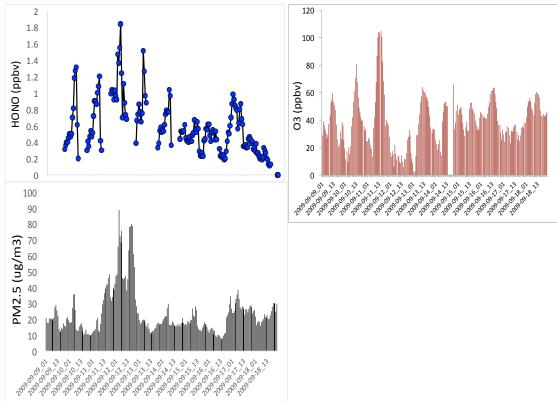
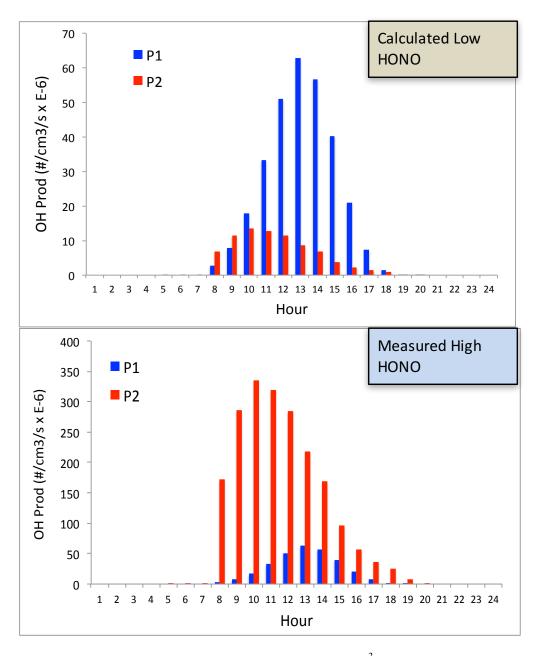
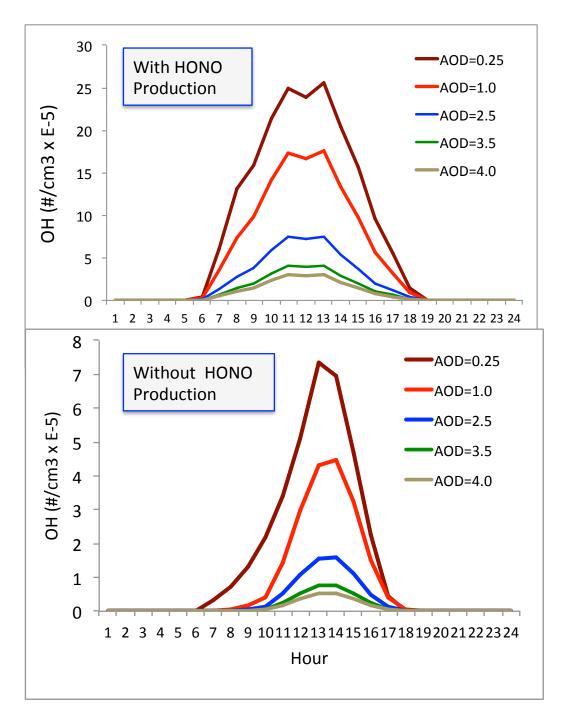


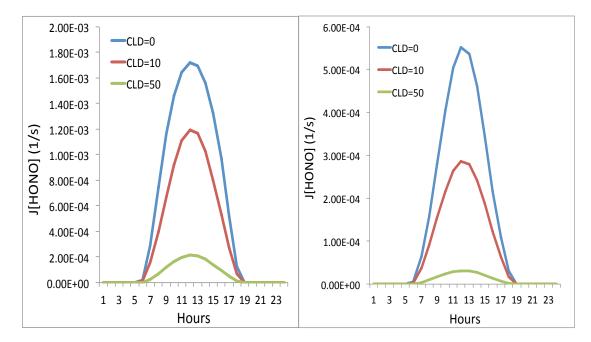
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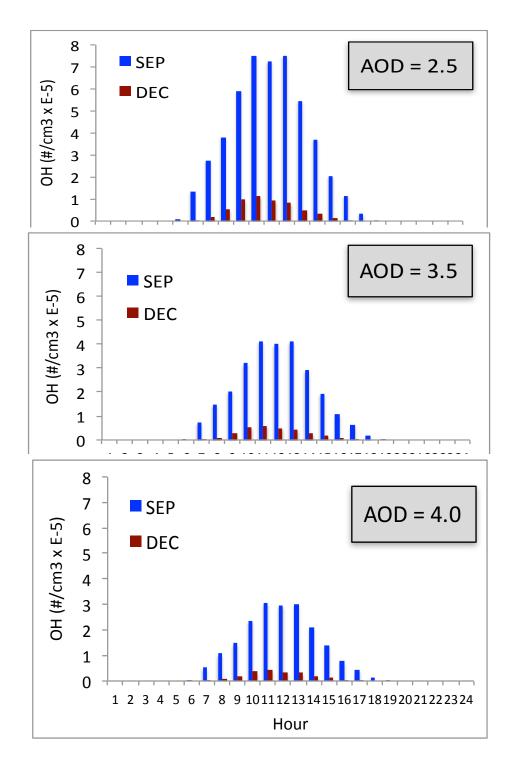
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