



2005–2017 ozone trends and potential benefits of local measures as deduced from air quality measurements in the north of the Barcelona Metropolitan Area

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Abstract

We analyzed 2005–2017 data sets on ozone (O_3) concentrations in an area (the Vic Plain) frequently affected by the atmospheric plume northward transport of Barcelona Metropolitan Area (BMA), the atmospheric basin of Spain recording the highest number of exceedances of the hourly O_3 information threshold ($180 \mu g m^{-3}$). We aimed at evaluating the potential benefits of implementing local-BMA short-term measures to abate emissions of precursors. To this end, we analyzed in detail spatial and time (interannual, weekly, daily and hourly) variations of concentration of O_3 and nitrogen oxides (NO and NO_2 , including remote sensing data for the latter) in April–September, and built a conceptual model for the occurrence of high O_3 episodes. Subsequently, a sensitivity analysis is done with the air quality (AQ) data to evaluate potential O_3 reductions in the North of the BMA on Sundays, compared with weekdays as a consequence of the reduction in regional emissions of precursors.

The results showed a generalized decreasing trend for regional background O_3 range (average $-1.4\% year^{-1}$), as well as the well-known increase of urban O_3 ($+1.8\% year^{-1}$), and higher urban NO decreasing slopes compared with those of NO_2 . The most intensive O_3 episodes in the Vic Plain are caused by (i) a relatively high regional background O_3 (due to a mix of continental, hemispheric–tropospheric and stratospheric contributions); (ii) intensive surface fumigation from mid-troposphere high O_3 upper layers arising from the concatenation of the vertical recirculation of air masses, but also by (iii) an important O_3 contribution from the northward transport/channeling of the pollution plume from the BMA.

The relevance of the local-daily O_3 contribution is supported by (i) the 5 times larger increase of the hourly exceedances of the O_3 information threshold on weekdays compared with Sundays; (ii) the occurrence of a marked O_3 Sunday decrease of the average diurnal concentrations over the Vic Plain; and (iii) a marked increase in concentrations of NO_2 (OMI- NO_2 remote sensing) over the Vic Plain–BMA region during days with the highest diurnal O_3 concentrations compared with the lowest.

We calculated the difference between the 75th percentile of O_x ($O_3 + NO_2$) diurnal concentrations recorded at one of the Vic Plain AQ monitoring stations for Wednesdays minus those of the 25th percentile of O_x for Sundays, equivalent to 1–2 days of emissions reductions in the BMA. A maximum decrease potential (by applying short-term measures to abate emissions of O_3 precursors) of $49 \mu g O_3 m^{-3}$ (32%) of the average diurnal concentrations was determined. Obviously, structurally implemented measures, instead of episodically, would result probably in important (and larger) additional O_x and O_3 decreases because not only the local O_3 coming from the BMA plume would be reduced but also the recirculated O_3 and thus the intensity of O_3 fumigation in the Plain. Therefore, it is highly probable that both structural and episodic measures to abate NO_x and volatile organic compounds (VOCs) emissions in the BMA would result in evident reductions of O_3 in the Vic Plain.



43 **Keywords:** tropospheric ozone, regional pollution, photochemistry, air quality trends.

44 1. Introduction

45 Tropospheric ozone (O_3) is a secondary atmospheric pollutant produced by the photooxidation of volatile
46 organic compounds (VOCs) in the presence of nitrogen oxides ($NO_x = NO + NO_2$). Its generation is enhanced
47 under high temperature and solar radiation (Monks et al., 2015 and references therein). Thus, O_3 maxima
48 occur generally in the afternoon, with the highest levels typically registered in summer, when exceedances of
49 regulatory thresholds are most frequent.

50 O_3 is one of the key air pollutants affecting human health and the environment (WHO, 2006, 2013a, 2013b;
51 GBD, 2016; Fowler et al., 2009; IPCC, 2013). According to EEA (2018), in the period 2013–2015, more than 95%
52 of the urban population in the EU-28 was exposed to O_3 levels exceeding the WHO guidelines set for the
53 protection of the human health (maximum daily 8-h average concentration of $100 \mu g m^{-3}$).

54 On a global scale, approximately 90% of the tropospheric O_3 is produced photochemically within the
55 troposphere (Stevenson et al., 2006; Young et al., 2013), the remaining part being transported from the
56 stratosphere (McLinden et al., 2000; Olson et al., 2001). The main global sink of tropospheric O_3 is photolysis
57 in the presence of water vapor. Dry deposition, mainly by vegetation, is also an important sink in the
58 continental planetary boundary layer (PBL) (Jacob and Winner, 2009).

59 On a regional scale, O_3 levels vary substantially depending on the different chemical environments within the
60 troposphere. O_3 chemical destruction is largest where water vapor concentrations are high, mainly in the lower
61 troposphere, and in polluted areas where there is direct O_3 destruction by titration. Thus, the hourly, daily and
62 annual variations in O_3 levels at a given location are determined by several factors, including the geographical
63 characteristics, the predominant meteorological conditions and the proximity to large sources of O_3 precursors
64 (Logan, 1985).

65 Southern Europe, especially the Mediterranean basin, is the most exposed to O_3 pollution in Europe (EEA,
66 2018) due to the specific prevailing meteorological conditions during warm seasons, regional pollutant
67 emissions, high biogenic VOCs' (BVOCs) emissions in spring and summer and the vertical recirculation of air
68 masses due to the particular orographic features that help stagnation–recirculation episodes (Millán et al.,
69 2000; EC, 2002, 2004; Millán, 2009; Diéguez et al., 2009, 2014; Valverde et al., 2016). Periods with high O_3
70 concentrations often last for several days and can be detected simultaneously in several countries. Lelieveld
71 et al. (2002) reported that during summer, O_3 concentrations are 2.5–3 times higher than in the hemispheric
72 background troposphere. High O_3 levels are common in the area, not only at the surface but also throughout
73 the PBL (Millán et al., 1997; Gangoiti et al., 2001; Kalabokas et al., 2007). Photochemical O_3 production is
74 favored due to frequent anticyclonic conditions with clear skies during summer, causing high insolation and
75 temperatures and low rainfall. Besides, the emissions from the sources located around the basin, which is
76 highly populated and industrialized, and the long-range transport of O_3 contribute to the high concentrations
77 (Millán et al., 2000; Lelieveld et al., 2002; Gerasopoulos, 2005; Safieddine et al., 2014).

78 In this context, the design of efficient O_3 abatement policies is difficult due to the following circumstances:

- 79 • The meteorology driving O_3 dynamics is highly influenced by the complex topography surrounding the
80 basin (see the above references for vertical recirculation of air masses and Mantilla et al., 1997;
81 Salvador et al., 1997; Jiménez and Baldasano, 2004; Stein et al., 2004).
- 82 • The complex nonlinear chemical reactions between NO_x and VOCs (Finlayson-Pitts and Pitts, 1993;
83 Pusede et al., 2015), in addition to the vast variety of the VOCs precursors involved and the
84 involvement of BVOCs in O_3 formation and destruction (Hewitt et al., 2011).
- 85 • The transboundary transport of air masses containing significant concentrations of O_3 and its
86 precursors, which contribute to increased O_3 levels, mainly background concentrations (UNECE, 2010).
- 87 • The contribution from stratospheric intrusions (Kalabokas et al., 2007).



- The fact that O_3 concentrations tend to be higher in rural areas (EEA, 2018), where local mitigation plans are frequently inefficient, because the emission of precursors takes place mostly in distant urban and industrial agglomerations.

Sicard et al. (2013) analyzed O_3 time trends during 2000–2010 in the Mediterranean and observed a slight decrease of annual O_3 averages (-0.4% per year) at rural sites, and an increase at urban and suburban stations ($+0.6\%$ and $+0.4\%$, respectively). They attributed the reduction at rural sites to the abatement of NO_x and VOCs emissions in the EU. Paradoxically, this led to an increase in O_3 at urban sites due to a reduction in the titration by NO. Their results also suggested a tendency to converge at remote and urban sites. Paoletti et al. (2014) also reported convergence in the EU and the US in the period 1990–2010 but found increasing annual averages at both rural and urban sites, with a faster increase in urban areas. Querol et al. (2016) determined that O_3 levels in Spain remained constant at rural sites and increased at urban sites in the period 2000–2015. This was suggested to be a result of the preferential reduction of NO versus NO_2 , supported by the lack of a clear trend in O_x ($O_3 + NO_2$). They also found that the target value was constantly exceeded in large areas of the Spanish territory, while most of the exceedances of the information threshold took place in July, mainly downwind of urban areas and industrial sites, and were highly influenced by summer heatwaves. The Vic Plain (located north of Barcelona) was the area registering the most annual exceedances of the information threshold in Spain, with an average of 15 exceedances per year per site.

In this study, we analyze NO, NO_2 and O_3 surface data around the Barcelona Metropolitan Area (BMA) and the Vic Plain, as well as NO_2 satellite observations, in the period 2005–2017, with the aim of better understanding the occurrence of high O_3 episodes in the area on a long-term basis. Previous studies in this region focused on specific episodes, whereas we aim at assessing the spatial distribution, time trends and temporal patterns of O_3 and its precursors, and the exceedances of the information threshold on a long time series. After understanding the 2005–2017 O_3 episodes, we aim to evaluate, as a first approximation using experimental data, the effect that episodic mitigation measures of O_3 precursors would have in the O_x concentrations in the Vic Plain.

2. Methodology

2.1. The area of study

The study is set in central Catalonia (Spain), in the north-eastern corner of the Iberian Peninsula (Figure 1). Characterized by a Mediterranean climate, summers are hot and dry with clear skies. In the 21st century, heat waves have occurred frequently in the area, often associated with high O_3 levels (Vautard et al., 2007; Guerova et al., 2007; Querol et al., 2016; Guo et al., 2017).

The capital city, Barcelona, is located on the shoreline of the Mediterranean Sea. Two sets of mountain chains lie parallel to the coastline (SW–NE orientation) and enclose the Pre-coastal Depression: the Coastal (250–500 m above sea level (a.s.l.)) and the Pre-Coastal (1000–1500 m a.s.l.) mountain ranges. The Vic Plain, situated 45–70 km North of Barcelona (500 m a.s.l.) is a 230 km² plateau that stretches along a S–N direction and is surrounded by high mountains (over 1000 m a.s.l.). The complex topography of the area protects it from Atlantic advections and continental air masses but also hinders the dispersion of pollutants (Baldasano et al., 1994). The two main rivers in the area (Llobregat and Besòs) flow perpendicularly to the sea and frame the city of Barcelona. Both rivers' valleys play an important role in the creation of air-flow patterns. The Congost River is a tributary to the Besòs River and its valley connects the Vic Plain with the Pre-coastal Depression.

The BMA stretches across the Pre-Coastal and Coastal Depressions and is a densely populated (>1500 people per km², MFom, 2017) and highly industrialized area with large emissions originating from road traffic, aircraft, shipping, industries, biomass burning, power generation and livestock.

During summer, the coupling of daily upslope winds and sea breezes may cause the penetration of polluted air masses up to 160 km inland, channeled from the BMA northward by the complex orography of the area. These air masses are injected at high altitudes (2000–3000 m a.s.l.) by the Pyrenean mountain ranges. At night



time, the land breeze prevails, and winds flow toward the sea followed by subsidence sinking of the air mass, which can be transported again by the sea breeze of the following day (Millán et al., 1997, 2000, 2002; Toll and Baldasano, 2000; Gangoiti, 2001; Gonçalves et al., 2009; Millán, 2014; Valverde et al., 2016). Under conditions of a lack of large-scale forcing and the development of a thermal low over the Iberian Peninsula that forces the confluence of surface winds from coastal areas toward the central plateau, this vertical recirculation of the air masses results in regional summer O₃ episodes in the Western Mediterranean. In addition, there might be external O₃ contributions, such as hemispheric transport or stratospheric intrusions (Kalabokas et al., 2007, 2008, 2017; Querol et al., 2017, 2018).

2.2. Air quality and meteorological data

We evaluated O₃ and NO_x AQ data together with meteorological variables and satellite observations of background NO₂.

The regional government of Catalonia (Generalitat de Catalunya, GC) has a monitoring network of stations that provides average hourly data of air pollutants (XVPCA, GC, 2017a, b). We selected a total of 25 stations (see Figure 2). Stations marked in green in Figure 2 met the following constraints: (i) location along the S–N axis (Barcelona–Vic Plain–Pre-Pyrenean Range); (ii) the station measures O₃; (iii) at least 9 years of data are available at the station in the period 2005–2017, with at least 75% data coverage from April to September. The remaining selected stations met the following criteria: (i) enough spatial and typology representativeness across the territory and (ii) availability of a minimum of 5 years of valid O₃ data in the period 2005–2017.

In addition, we selected wind and temperature data from five meteorological stations from the Network of Automatic Meteorological Stations (XEMA, Meteocat, 2017) closely located to the previously selected AQ stations, as well as solar radiation data from two solar radiation sites from the Catalan Network of Solar Radiation Measurement Stations (ICAEN-UPC, 2018) located in the cities of Girona and Barcelona.

2.3. Data analysis

2.3.1. O_x calculations

We calculated O_x concentrations to better interpret O₃ dynamics. Kley and Gleiss (1994) proposed the concept of O_x to improve the spatial and temporal variability analysis by decreasing the effect of titration of O₃ by NO with the subsequent consumption of O₃ in areas where NO concentrations are high. Concentrations were transformed to ppb units using the conversion factors at 20 °C and 1 atm (DEFRA, 2014).

O_x concentrations were only calculated if there were at least six simultaneous hourly recordings of O₃ and NO₂ from 12:00 to 19:00 h, June–August, in the period 2005–2017. The stations used for these calculations were those located along the S–N axis (Barcelona–Vic Plain–Pre-Pyrenean Range).

2.3.2. Spatial variation

To study the spatial distributions of NO, NO₂, O₃ and O_x across the region we calculated June–August averages (months recording the highest concentrations of O₃ in the area) from hourly concentrations provided by all the selected AQ sites. For each of them, we calculated daily averages and daytime high averages (12:00 to 19:00 h).

2.3.3. Time trends

By means of the Mann–Kendall method, we analyzed time trends for NO, NO₂ and O₃ for the period 2005–2017. In addition, we used the Theil–Sen statistical estimator (Theil, 1950; Sen, 1968) implemented in the R package Openair (Carslaw and Ropkins, 2012) to obtain the regression parameters of the trends (slope, uncertainty and p-value) estimated via bootstrap resampling. We examined the annual time trends of seasonal averages (April–September) for each pollutant. Data used for these calculations were selected according to the recommendations in EMEP-CCC (2016): the stations considered have at least 10 years of data (75% of the total period considered, 2005–2017) and at least 75% of the data is available within each season. In addition, we analyzed annual time trends of tropospheric NO₂ measured by satellite along the S–N axis and of



182 greenhouse gases (GHGs) emitted in Catalonia and the average number of vehicles entering the city of
 183 Barcelona.

184 **2.3.4. Assessment of O₃ objectives according to air quality standards**

185 We identified the maximum daily 8-hour average concentrations by examining 8-h running averages using
 186 hourly data in the period 2005–2017. Each 8-h average was assigned to the day on which it ended (i.e., the
 187 first average of one day starts at 17:00 h on the previous day), as determined by EC (2008).

188 To assess the time trends and patterns of the Exceedances of Hourly Information Thresholds (EHITs)
 189 established by EC (2008) (hourly mean of O₃ concentration greater than 180 µg m⁻³), we used all the data,
 190 independently of the percentage of data availability.

191 **2.3.5. Tropospheric NO₂ column**

192 We used daily tropospheric NO₂ column satellite measurements using the Ozone Monitoring Instrument (OMI)
 193 spectrometer aboard NASA's Earth Observing System (EOS) Aura satellite (see OMI, 2012; Krotkov and
 194 Veefkind, 2016). The measurements are suitable for all atmospheric conditions and for sky conditions where
 195 cloud fraction is less than 30% binned and averaged into 0.25° × 0.25° global grids. We then analyzed daily
 196 average Tropospheric Column NO₂ measurements from 2005 to 2017 aiming at two different goals. On the
 197 one hand, to quantify the tropospheric NO₂ in the area along the S–N axis and obtain annual time trends and
 198 monthly/weekly patterns. On the other hand, to assess qualitatively the tropospheric NO₂ across a regional
 199 scale (Western Mediterranean Europe) in two different scenarios, by means of visually finding patterns that
 200 might provide a better understanding of O₃ dynamics in our area of study. The scenarios were: days with the
 201 maximum 8-h O₃ average above the 75th percentile at the Vic Plain stations, and days with the maximum
 202 below the 25th percentile. See selected regions for retrieval of NO₂ satellite measurements in Figure S1.

203 **2.3.6. Time conventions**

204 When expressing average concentrations, the times shown indicate the start time of the average. For example,
 205 12:00–19:00 h averages take into account data registered from 12:00 h to 19:59 h. All times are expressed as
 206 local time (UTC + 1 hour during winter and UTC + 2 hours during summer) and the 24-hour time clock
 207 convention is used.

208 **3. Results and discussion**

209 **3.1. Spatial variation of concentrations**

211 We analyzed the mean NO, NO₂, O₃ and O_x concentrations (June to August) in the study area in the period
 212 2005–2017.

213 As expected, the highest NO and NO₂ concentrations are registered in urban/suburban (U/SU) traffic sites in
 214 and around Barcelona (MON, GRA, MNR and CTL, 7–10 µg NO m⁻³ and CTL and MON 30–36 µg NO₂ m⁻³). Also,
 215 as expected, the remote high-altitude rural background (RB) sites (MSY and MSC) register the lowest NO (<1
 216 µg m⁻³) and NO₂ (2–4 µg m⁻³) concentrations, see Figure S2.

217 The lowest June–August average O₃ concentrations (45–60 µg m⁻³) are recorded in the same U/SU traffic sites
 218 (MON, GRA, MNR and CTL) where titration by NO is notable, while the highest ones (>85 µg m⁻³) are recorded
 219 at the RB sites, MSC being the station recording the highest June–August O₃ levels (102 µg m⁻³). These spatial
 220 patterns are significantly different when we consider the 8-h daily averages of O₃ concentrations for June–
 221 August 12:00–19:00 h (Figure 3a). Thus, these concentrations are repeatedly high (85–115 µg m⁻³) in the whole
 222 area of study. The highest O₃ concentrations (>107 µg m⁻³) were recorded at the four sites located downwind
 223 of BMA along the S–N corridor (MSY, TON, VIC and MAN), and downwind of Tarragona (PON, RB station).
 224 Figure 3a also shows a positive O₃ gradient along the S–N axis (O₃ levels increase farther north) following the
 225 BMA plume transport and probably an increase of the mixing layer height (MLH). The higher O₃ production
 226 and/or fumigation in the northern areas are further supported by the parallel northward increasing O_x gradient



(δO_x , Figure 3b). Time series show that in 85% of the valid data in June–August (849 out of 1001 days in 2005–2017) this positive gradient is evident between CTL and TON ($\delta O_{x\text{ TON-CTL}} > 0$). The average O_x increase between CTL in Barcelona and TON is 15 ppb. Taking into account the low NO_2 concentrations registered at this station, this is equivalent to approximately $29 \mu\text{g m}^{-3}$ of O_3 (+30% O_x in TON compared with CTL).

Thus, TON at the Vic Plain records the highest 12:00–19:00 h, June–August O_x and O_3 concentrations in the study area. The MNR site also exhibits very high O_x levels (Figure 3b) but these are mainly caused by primary NO_2 associated with traffic emissions.

3.2. Time patterns

3.2.1. Annual trends

Table 1 shows the results of the trend analysis of NO , NO_2 , O_3 and O_x averages (April to September, the O_3 season according to the European AQ Directive) by means of the Mann–Kendall test.

NO_x levels exhibit a generalized and progressive decrease during the time period across Catalonia. In particular, NO_2 tended to decrease along the S–N axis during the period (U/SU sites CTL, MON and MAN registered -1.6 , -2.0 and -1.3% year^{-1} , respectively, with statistical significance in all cases). A similar trend was found for NO in these stations, with higher negative slopes (-2.2 , -4.3 and -1.1% year^{-1} , the latter without statistical significance).

The annual averages of tropospheric NO_2 across the S–N axis decreased by 35% from 2005 to 2017 (-3.4% year^{-1} with statistical significance). The marked drop of NO_2 from 2007–2008 can be attributed to the reduction of emissions associated with the financial crisis starting in 2008. The time trends of average traffic (number of vehicles) entering Barcelona City on working days from 2005 to 2016 (Ajuntament de Barcelona, 2010, 2017) and the GHGs emitted in Catalonia attributed to industry and power generation sectors calculated from the Emissions Inventories published by the Regional Government of Catalonia from 2005 to 2016 (GC, 2017c) (Figure 4a) support this hypothesis. We found both decreasing trends to be statistically significant but the GHG emissions decreasing trend is significantly higher (-3.8% year^{-1}) than the traffic (-1.2% year^{-1}), which suggests that the crisis had a more severe effect on industry and power generation than on road traffic. This is also supported by a larger decrease of GHG emissions and OMI- NO_2 from 2005–2007 (precrisis) to 2008 (start of the crisis) than BMA traffic counting and urban NO_x levels (without a 2007–2008 steep change and a more progressive decrease, Figure 4b). Thus, in the BMA, the financial crisis caused a more progressive decrease (without a 2007–2008 steep change) of the circulating vehicles and therefore its associated emissions.

April–September O_3 and O_x mean concentration trends are shown in Table 1. The data show that seven out of the eight RB sites registered slight decreases in O_3 concentrations during the period (BdC, AGU and STP; -1.6% year^{-1} , -1.1% year^{-1} and -1.4% year^{-1} , respectively, in all cases with statistical significance) while in BEG, PON, LSE and GAN the trends were not significant. As in several regions of Spain and Europe (Sicard et al., 2013; Paoletti et al., 2014; EEA, 2016; Querol et al., 2016; EMEP, 2016), the opposite trends are found for U/SU sites, with increases in O_3 concentrations during the period at some stations (CTL, MON, MAN, MAT, MNR and ALC; $+0.4$ to $+3.2\%$ year^{-1} all with statistical significance). When considering O_x , the increasing trends in U/SU sites are neutralized in some cases (CTL, MON, MAN, MAT and ALC). This, and the higher NO decreasing slopes compared with those of NO_2 , support the hypothesis that the U/SU O_3 increasing trends are probably caused by less O_3 titration (due to decrements in NO levels) instead of a higher O_3 generation.

3.2.2. Monthly and daily patterns

Figure 5a shows 2005–2017 monthly average hourly O_3 concentrations measured at sites along the S–N axis, showing the occurrence of chronic-type episodes with repeated high O_3 concentrations (90 – $135 \mu\text{g m}^{-3}$) in the



271 afternoon of April–September days at the Vic Plain sites (TON, VIC, MAN) and the remote RB sites (MSY and
 272 PAR).

273 Typically, at the remote RB stations, O₃ concentrations are high during the whole day throughout the year,
 274 and daily O₃ variations are narrower than at the other stations, with high average levels even during October–
 275 February (MSY: 50–70 and PAR: 50–80 µg m⁻³). During the night these mountain sites are less affected by NO
 276 titration, leading to high daily O₃ average concentrations. However, in summer, midday–afternoon
 277 concentrations are relatively lower than at the stations located in the S–N valley (TON, VIC, MAN).

278 Regarding monthly average daily O_x (Figure 5b), the profiles of RB sites TON and MSY are very similar to the
 279 respective O₃ profiles. In the case of the BMA U/SU sites (CTL, MON, GRA), the nocturnal O_x concentrations
 280 increase with respect to O₃ due to the addition of secondary NO₂ from titration. Midday–afternoon O_x levels
 281 are much lower at the BMA U/SU stations than those in the S–N valley (MAN, TON), similarly to O₃ levels,
 282 supporting the contribution of local-regional O₃ from the BMA plume and/or from the fumigation of high-
 283 altitude reserve strata as MLH grows (Millán et al., 1997, 2000; Gangoiti et al., 2001; Querol et al., 2017).

284 3.2.3. Weekly patterns

285 As stated before, 8-h day maxima of O₃ concentrations are generally recorded at 12:00–19:00 h. Accordingly,
 286 Figure 6 shows the O₃ weekly patterns for these O₃ average concentrations. As expected, the variation of intra-
 287 annual concentration values is pronounced in the Vic Plain sites (TON, VIC, MAN; 20–45 µg m⁻³ in December–
 288 January versus 110–125 µg m⁻³ in July), due to the higher summer photochemistry, the more frequent summer
 289 BMA plume transport (due to intense sea breezing) and fumigation from upper atmospheric reservoirs across
 290 the S–N axis, and of the high O₃ titration in the populated valleys in winter. However, at the remote mountain
 291 sites of MSY and PAR, the intra-annual variability is much reduced (70–80 µg m⁻³ in December versus 100–120
 292 µg m⁻³ in July) probably due to the reduced effect of NO titration at these higher altitude sites, and the
 293 influence of high-altitude O₃ regional reservoirs.

294 During the year, CTL, MON and GRA (U/SU sites around BMA) register very similar weekly patterns of the 8-h
 295 maxima, with a marked and typical high O₃ weekend effect, i.e., higher O₃ levels than during the week due to
 296 lower NO concentrations. From April to September, CTL O₃ 8-h concentrations are lower than MON's and
 297 GRA's (the latter located to the north of BMA following the sea breeze air mass transport), despite being very
 298 similar from October to March (when sea breezes are weaker). An O₃ weekend effect is also clearly evident
 299 during the winter months in the Vic Plain sites (TON, VIC, MAN) and MSY. However, from June to August, a
 300 marked inverse weekend effect is clearly evident at these same sites, with higher O₃ levels during weekdays.
 301 This points again to the clear influence of the emission of precursors from the BMA on the O₃ concentrations
 302 recorded at these inland sites.

303 3.3. Peak O₃ concentrations patterns along the S–N axis

304 July is the month of the year when most of the annual exceedances of the O₃ EHITs are recorded in Spain
 305 (Querol et al., 2016), including our area of study. Figure 7 shows the average O₃ and O_x July hourly
 306 concentrations along the S–N axis during 2005–2017. A progressive time-shift and a marked positive
 307 northward gradient of O₃ and O_x maxima are shown, pointing again to the gradual increase of O₃ and O_x due
 308 to the plume transport and fumigation from upper reservoirs as MLH grows.

309 Figure 8a shows the 2005–2017 trends of the EHITs from the European AQ Directive (>180 µg m⁻³ h⁻¹ mean;
 310 EC, 2008) registered at the selected sites in the S–N valley, as well as the average temperatures measured
 311 during July at early afternoon near Vic (at Gurb meteorological site), the background NO₂ measured by OMI
 312 (June to August) and the average solar radiation measured in Girona and Barcelona (June to August). In 2005,
 313 2006, 2010, 2013, 2015 and 2017, the highest EHITs at almost all the sites were recorded. Temperature and
 314 insolation seem to have a major role in the occurrence of EHITs in 2006, 2010, 2015 and 2017. The effect of
 315 heat waves on O₃ episodes is widely known (Solberg et al., 2008; Meehl et al., 2018; Pyrgou et al., 2018; among
 316 others). However, because the emissions of precursors have clearly decreased (–30% decrease on June to



August OMI-NO₂ levels across the S–N axis from 2005 to 2017; -2.7% year⁻¹ with statistical significance) the number of EHITs recorded in the warmest years has probably decreased with respect to a scenario where emissions would have been maintained. In any case, some years (for example 2009 and 2016) seem to be out of line for temperature and insolation being the driving forces, and other major causes also have to be relevant, with further research needed to interpret fully interannual trends. Otero et al. (2016) found that temperature is not the main driver of O₃ in the South-western Mediterranean, as it is in Central Europe, but the O₃ levels recorded the day before (a statistical proxy for the occurrence of Millán et al. (1997)'s vertical recirculation of air masses). Again, the Vic Plain sites (TON, VIC, MAN) recorded most (75%) of the EHITs reported by the AQ monitoring stations in Catalonia (25%, 34% and 16%, respectively). The higher urban pattern of MAN, as shown by the higher NO concentrations, with respect to TON, might account for both the lower exceedances and the different interannual patterns.

Figure 8b shows that most EHITs occurred in June and July (30% and 57%, respectively), with much less frequency in May, August and September (6%, 8% and <1%, respectively). Although temperatures are higher in August than in June, the latter registers significantly more EHITs, probably due to both the stronger solar radiation and the higher concentrations of precursors (such as NO₂, see OMI-NO₂ and solar radiation in Figure 8b).

Figure 8c shows that EHITs occurred mainly between Tuesday and Friday (average of 19% of occurrences per day). On weekends and Mondays, EHITs were clearly lower (average of 9% of occurrences per day) than during the rest of the week, probably due to: (i) the lower emissions of anthropogenic O₃ precursors (such as NO_x, see OMI-NO₂) during weekends and (ii) to the effect of the lower Sunday emissions in the case of the lower exceedances recorded during Mondays. During weekends and in August, OMI-NO₂ along the S–N axis is relatively lower (-29% weekday average and -43% in August with respect to March) following the emissions patterns associated with industrial and traffic activity that drop during vacations and weekends (Figure 8). NO_x data from AQ monitoring sites follow similar patterns (not shown here).

Figure 8d shows that the frequency of occurrence of the EHITs at MSY (45 km north of Barcelona) is lower and earlier (maxima at 14:00 h) than at Vic Plain sites (TON, VIC, MAN). The EHITs occurred mostly at 15:00, 16:00, 16:00 and 19:00 h at TON, VIC, MAN and PAR (53, 63, 72 and 105 km north of Barcelona), respectively. PAR registered not only much later EHITs, but a much lower number than TON-VIC-MAN sites, again confirming the progressive O₃ maxima time-shift northward of Barcelona.

The results in Figure 9 clearly show that during non-EHIT days, the daily O₃ patterns are governed by the morning–midday concentration growth driven to fumigation and photochemical production, while on EHIT days there is a later abrupt increase, with maxima being delayed as we increase the distance from Barcelona along the S–N axis. This maximal second increase of O₃ is clearly attributable to the influence of the transport of the plume of the BMA (horizontal transport), as the secondary NO₂ peak at 15:00 h (Figure 9 left bottom), and the wind patterns (see Figure S3) seem to support. The differences in the late hourly O₃ concentration increases in EHIT versus non-EHIT days are even more evident when calculating hourly O₃ slopes (hourly increments or decrements of concentrations), Figure 9 (right). The first increment (fumigation and photochemistry) makes O₃ levels scale up to 120 $\mu\text{g m}^{-3}$ during EHIT episodes and to nearly 100 $\mu\text{g m}^{-3}$ during non-EHIT days. In EHIT days, the later peak (transport from BMA and causing most of the 180 $\mu\text{g m}^{-3}$ exceedances) in the O₃ slope occurs again between 14:00 h and 20:00 h, depending on the distance to BMA, but this feature is not observed on non-EHIT days.

3.4. A conceptual model for O₃ episodes in NE Spain

Figure 10 depicts the basic atmospheric dynamics in the study area during a typical summer day, when the atmospheric conditions are dominated by mesoscale circulations. According to the previous references, indicated in Figure 10 with enclosed numbering (coinciding with the numbering below) the following O₃ contributions to surface concentrations in the study area can be differentiated:

- a. Vertical recirculation of O₃-rich air masses, which create reservoir layers of aged pollutants.



- 364 b. Vertical fumigation of O_3 from the above reservoirs and the following sources aloft if the MLH growth is
365 large enough:
366 b.1. Regional external O_3 layers (from other regions of southern Europe, such as southern France, Italy,
367 Portugal and Tarragona).
368 b.2. High free tropospheric O_3 background due to hemispheric long-range transport.
369 b.3. High free tropospheric O_3 background due to stratospheric intrusions.
370 c. Horizontal transport of O_3 . Diurnal BMA plume northward transported and channeled into the Besòs–
371 Congost valleys.
372 d. Local production of O_3 from precursors.

373 During summer, the intense land heating due to strong solar radiation begins early in the morning. The
374 associated convective activity produces morning fumigation processes (b in Figure 10) that bring down O_3 from
375 the reservoir layers aloft, creating sharp increases in O_3 concentrations in the morning (see Figures 9 and S3).
376 The breeze transports air masses from the sea inland and creates a compensatory subsidence of aged
377 pollutants (including O_3) previously retained in reservoir and external layers and high free troposphere
378 background aloft (Millán et al., 1997, 2000; Gangoiti et al., 2001). This subsided O_3 then affects the marine
379 boundary layer and reaches the city the following day with the sea breeze, producing nearly constant O_3
380 concentrations in the city during the day (Figure S3 and Figure 7). As the breeze develops, coastal emissions
381 and their photochemical products are transported inland, generating the BMA plume (c in Figure 10) that, in
382 addition to the daily generated O_3 , also contains recirculated O_3 from the marine air masses. Furthermore,
383 during the transport to the Vic Plain, new O_3 is produced (d in Figure 10) by the intense solar radiation and the
384 O_3 precursors emitted along the way (e.g., BVOCs from vegetation, NO_x from industrial and urban areas,
385 hydrocarbons).

386 This new O_3 gets mixed with the BMA plume and channeled northward to the S–N valleys until it reaches the
387 Vic Plain and the southern slopes of the Pre-Pyrenees. As the BMA plume (loaded with O_3 and precursors)
388 travels northward, a second increase in O_3 concentrations can be observed in the daily cycles of O_3 at these
389 sites, (see Figures 9 and S3). This was described as the second O_3 peak by Millán et al. (2000).

390 The land use in the Vic Plain is mainly agricultural, with few forested and humid areas which favor heating,
391 diurnal dryness and high temperatures during summer. This fact may enhance the growth of the MLH (up to
392 3000 m a.s.l.; Querol et al., 2017) compared with coastal areas close to Barcelona, where the fresh sea breeze
393 fluxes account for a lower MLH (Soriano et al., 2001). This marked MLH increase may produce a preferential
394 and intensive top-down O_3 transport (b in Figure 10) from upper O_3 layers (a, b.1, b.2 and b.3 in Figure 10),
395 contributing to high O_3 surface concentrations in the Vic Plain. During the sea/mountain breezes'
396 development, some air masses are injected upward to the N and NW return flows (controlled by the synoptic
397 circulations dominated by the high-pressure system over the Azores) aloft helped by the orography (e.g.,
398 southern slopes of mountains) and again transported back to the coastal areas where, at late evening/night it
399 can accumulate at certain altitudes in stably stratified layers.

400 Later, at night, land breezes returning to the coastal areas develop. Depending on the orography, these
401 drainage flows of colder air traveling to the coastal areas can accumulate on the surface or keep flowing to
402 the sea. The transported O_3 is consumed along the course of the drainage flows by deposition and titration.
403 Next day, the cycle starts anew, producing almost closed loops enhancing O_3 concentrations throughout the
404 days in the area. When the loop is active for several days, multiple O_3 EHITs occur over the Vic Plain.

405 The main complexity of this system arises from the fact that all these vertical/horizontal,
406 local/regional/hemispheric/stratospheric contributions are mixed and all contribute to surface O_3
407 concentrations with different proportions that may largely vary with time and space across the study area.
408 However, for the most intense O_3 episodes, the local-regional contribution might be very relevant to cause
409 EHITs in the region. Furthermore, the intensity and frequency of O_3 episodes are partially driven by the
410 occurrence of heat waves in summer and spring. If local and regional emissions of precursors are high, the



intensity of the episodes will also be high. Thus, even though heat wave occurrences increase the severity of O_3 episodes, an effort to reduce precursors should be undertaken to decrease their intensity.

The conceptual model that we propose for the generation of the O_3 episodes in 2005–2017 for the S–N corridor BMA–Vic Plain–Pre-Pyrenees does not differ very much from the one proposed by Millán et al. (1997, 2000, 2002), Gangoiti et al. (2001), Kalabokas et al. (2007, 2008, 2017), Millán (2014) and Querol et al. (2018) for other regions of the Mediterranean basin, including Spain, or described in the same area for specific episodes (Toll and Baldasano, 2000; Gonçalves et al., 2009; Valverde et al., 2016; Querol et al., 2017). However, in the present model the role of the local-regional emissions on the occurrence of O_3 EHITs is clearly more relevant. Thus, our results demonstrate an increase in the EHITs northward from Barcelona to around 70 km and a decrease from there to 100 km from Barcelona following the same direction. There is also a higher frequency of occurrence of these in July (and June) and from Tuesday to Friday and a time-shift of the frequency of occurrence of EHITs from 45 to 100 km. The mountain site of MSY (located at 700 m a.s.l.) registered many fewer EHITs than the sites in the valleys (TON–VIC–MAN, 460–600 m a.s.l.) during the period, showing the key role of the valley channeling of the high O_3 and precursors BMA plume in July (when sea breeze and insolation are more intense). Furthermore, at the Vic Plain, we detected an inverse O_3 weekend effect, suggesting that local/regional anthropogenic emissions of precursors play a key role in increasing the number of EHITs on working days, with a Friday/Sunday rate of 5 for VIC for 2005–2017. Despite this clear influence of the BMA plume on EHITs' occurrence, Querol et al. (2017) demonstrated that at high atmospheric altitudes (2000–3000 m a.s.l.) high O_3 concentrations are recorded, in many cases reaching $150 \mu\text{g m}^{-3}$ due to the frequent occurrence of reservoir strata. As also described above, the higher growth of the MLH in TON–VIC–MAN as compared with the coastal area accounts also for higher top-down O_3 contributions. On the other side, close to the Pyrenees (PAR station), large forested and more humid areas give rise to a thinner MLH, hindering O_3 fumigation too. Furthermore, in these more distant northern regions O_3 consumption by ozonolysis of BVOCs might prevail over production due to weaker solar radiation during the later afternoon.

Figure 11 shows the distribution of average background OMI- NO_2 levels across the Western Mediterranean Basin in two different scenarios: when the O_3 levels in the Vic Plain are low (left) or high (right). To this end, we averaged the values from VIC and TON (in the Vic Plain) from all the maximum daily 8-h mean O_3 concentrations calculated for all the days in July within 2005–2017, and we calculated the 25th (93 out of 370 days, $105 \mu\text{g m}^{-3}$) and 75th (93 days, $139.5 \mu\text{g m}^{-3}$) percentiles of all the data (P25 and P75, respectively). For both scenarios, NO_2 concentrations are highest around large urban and industrial areas, including Madrid, Porto, Lisbon, Barcelona, Valencia, Paris, Frankfurt, Marseille and especially the Po Valley. The shipping routes toward the Gibraltar Strait and around the Mediterranean can be observed, as well as important highways such as those connecting Barcelona to France and Lyon to Marseille. As expected, the mountain regions (the Pyrenees and the Alps) are the areas with lower NO_2 . Regional levels of background OMI- NO_2 in the P75 scenario are markedly higher with hotspots intensified and spanning over broader areas. Over Spain, new hotspots (marked in yellow), such as the coal-fired power plants in Asturias (a), ceramic industries in Castelló (c) and the coal-fired power plant in Andorra, Teruel (b), appear; in the latter case, with the pollution plume being channeled along the Ebro Valley with a NW transport. Furthermore, it is important to highlight that the maxima background NO_2 along the eastern coastline in Spain, including the BMA, tend to exhibit some north-northwest displacement, when compared with the P25 scenario, thus pointing to the relevance of the local emissions in causing inland O_3 episodes.

These qualitative results suggest in general less synoptic forcing in Western Europe in the P75 scenario; hence, in these conditions NO_2 is accumulated across the region and especially around its sources. In the east coast of the Iberian Peninsula, mesoscale circulations tend to dominate, hence the northwest displacement (taking the coastal regions as a reference) of the background NO_2 . The bottom part of Figure 11 zooms our study area and shows the maximum daily 8-h mean O_3 concentrations in all the selected AQ sites averaged for both scenarios. As shown in the P75 scenario, NO_2 is significantly intensified across Catalonia, especially north of the BMA spreading to the Vic Plain. Comparing O_3 in both scenarios, in the P75 the O_3 levels are much higher



459 (mostly $>105 \mu\text{g m}^{-3}$), across the region except the urban sites in Barcelona (due to NO titration), reaching up
460 to $154 \mu\text{g m}^{-3}$ in the Vic Plain.

461 Conversely, in the P25 scenario, background NO_2 concentrations are lower, the BMA NO_2 spot is significantly
462 smaller and spreads along the coastline rather than being displaced to the north-northwest. In this case,
463 synoptic flows seem to weaken sea breeze circulations and vertical recirculation, thus reducing the amount of
464 background NO_2 and the inland transport from the coast. In these conditions, O_3 levels are markedly lower
465 across the territory, the RB PON site (downwind of the city/industrial area of Tarragona) being the one
466 recording the maximum daily 8-h mean O_3 concentration ($99 \mu\text{g m}^{-3}$).

467 3.5. Sensitivity analysis for O_x using experimental data

468 We demonstrated above that the lower anthropogenic emissions of O_3 precursors in the BMA during
469 weekends cause lower O_3 and O_x levels in the Vic Plain than during working days (inverse O_3 weekend effect).
470 To apply a sensitivity analysis using experimental data for the O_3 levels in the Vic Plain if BMA's emissions were
471 reduced, we compared weekend O_3 and O_x patterns with weekdays considering only data from June and July
472 (August OMI- NO_2 levels are markedly lower, Figure 8b, therefore this month was not included).

473 Figure 12 shows the average O_x concentrations in TON according to the day of the week for the period
474 considered. Despite the large variability in extreme values (i.e., maximum values with respect to minimum
475 values, represented by whiskers), the interquartile range is quite constant on all the weekdays (between 13.6
476 and 17.3 ppb). The average O_x decrease between the day with highest O_x levels (Wednesday) and the day with
477 the lowest O_x levels (Sunday) is 6.5 ppb (approximately $13 \mu\text{g O}_3 \text{ m}^{-3}$, 10% decrease). Thus, we calculated the
478 difference between the P75 of O_x values observed on Wednesdays minus the P25 of O_x values on Sundays,
479 equivalent to 1–2 days of emissions reductions in the BMA. In this case, it is a feasible scenario to consider a
480 maximum decrease of 24.5 ppb (approximately $49 \mu\text{g O}_3 \text{ m}^{-3}$, 32% decrease) after 1–2 days of mitigation
481 measures of precursor emissions in the BMA. Obviously, if these mitigation measures would be implemented
482 structurally, instead of episodically, probably the O_x and O_3 decreases would be larger because not only the
483 local O_3 coming from the BMA plume would be reduced but also the recirculated O_3 and thus the intensity of
484 O_3 fumigation in the Plain. Therefore, it is probable that both structural and episodic measures to abate VOCs
485 and NO_x emissions in the BMA would result in evident reductions of O_3 in the Vic Plain, as evidenced by
486 modeling tools by Valverde et al. (2016).

487 4. Conclusions

488 We analyzed 2005–2017 data sets on ozone (O_3) concentrations in an area frequently affected by the
489 northward atmospheric plume transport of Barcelona Metropolitan Area (BMA) to the Vic Plain, the area of
490 Spain recording the highest number of exceedances of the hourly O_3 information threshold (EHIT, $180 \mu\text{g m}^{-3}$).
491 We aimed at evaluating the potential benefits of implementing local short-term measures to abate
492 emissions of precursors. To this end, we analyzed in detail spatial and time (interannual, weekly, daily and
493 hourly) variations of the concentration of O_3 and nitrogen oxides (including remote sensing data for the latter)
494 in April–September and built a conceptual model for the occurrence of high O_3 episodes. Finally, a sensitivity
495 analysis is done with the AQ data to evaluate potential O_3 reductions in the North of the BMA on Sundays,
496 compared with weekdays, as a consequence of the reduction of emissions of precursors.

497 Results showed a generalized decrease trend for regional background O_3 ranging from -1.1 to $-1.6\% \text{ year}^{-1}$,
498 as well as the well-known increase of urban O_3 ($+0.4$ to $+3.2\% \text{ year}^{-1}$) and higher urban NO decreasing slopes
499 than those of NO_2 (-2.2 to -4.3 and -1.3 to $-2.0\% \text{ year}^{-1}$, respectively), that might account in part for the
500 urban O_3 increase.

501 The conceptual model for the most intensive O_3 episodes in the North of the BMA is based on a relatively high
502 regional background O_3 (due to a mix of continental, hemispheric–tropospheric and stratospheric
503 contributions) on top of which intensive surface fumigation from the mid-troposphere high O_3 upper layers



504 arising from the concatenation of the vertical recirculation of air masses (as a result of the interaction of a
505 complex topography with intensive spring–summer sea and mountain breezes circulations (Millán et al., 1997,
506 2000; Gangoiti et al., 2001; Valverde et al., 2016; Querol et al., 2017). However, we noticed that for most EHIT
507 days in the Vic Plain, the exceedance occurs when an additional contribution is added to the previous two: O₃
508 supply by the channeling of the BMA pollution plume along the S–N valley connecting BMA and Vic. Thus,
509 despite the large external O₃ contributions, structural and short-time local measures to abate emissions of
510 precursors might clearly influence spring–summer O₃ in the Vic Plain. This is supported by (i) the reduced
511 hourly exceedances of the O₃ information threshold recorded on Sundays at the Vic AQ monitoring site (9 in
512 2005–2017) compared with those on Fridays (47), as well as by (ii) the occurrence of a typical and marked
513 Sunday O₃ pattern at the BMA AQ monitoring sites and an also marked but opposite one in the sites of the Vic
514 Plain; and (iii) marked increase of remote sensing OMI-NO₂ concentrations over the BMA and northern regions
515 during days of the P75 diurnal O₃ concentrations compared with those of the P25.

516 Finally, we calculated the difference between the P75 of O_x diurnal concentrations recorded at one of the Vic
517 Plain AQ monitoring stations for Wednesdays minus those of the P25 percentile of O_x for Sundays, equivalent
518 to 1–2 days of emissions reductions in the BMA. A maximum decrease potential by applying short-term
519 measures of 24.5 ppb (approximately 49 µg O₃ m⁻³, 32% decrease) of the diurnal concentrations was
520 calculated. Obviously, structurally implemented measures, instead of episodic ones, would result probably in
521 important additional O_x and O₃ abatements because not only the local O₃ coming from the BMA plume would
522 be reduced but also the recirculated O₃, and thus the intensity of O₃ fumigation on the Plain. Therefore, it is
523 highly probable that both structural and episodic measures to abate NO_x and VOCs emissions in the BMA
524 would result in evident reductions of O₃ in the Vic Plain.

525 **Author contributions**

526 JM performed the data compilation, treatment and analysis with the aid of XQ, CC and ME. JM, CC, ME, JB, AA
527 and XQ contributed to the discussion and interpretation of the results. JM and XQ wrote the manuscript. JM,
528 CC, ME, JB, AA and XQ commented on the manuscript.

529 **Competing interests**

530 The authors declare that they have no conflict of interest.

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785 **FIGURE CAPTIONS**

786 Figure 1. Location and main topographic features of the area of study.

787

788 Figure 2. Location (left) and main characteristics (right) of the selected air quality monitoring sites (S–N axis:
 789 green squares on the map and shaded gray on the table, rest of stations: white squares) and
 790 meteorological/solar radiation stations (red circles) selected for this study. Types of air quality monitoring sites
 791 are urban (traffic or background: UT, UB), suburban (traffic, industrial or background: SUT, SUI, SUB) and rural
 792 (background or industrial: RB, RI). PLR (Palau Reial air quality monitoring site) and BCN (Barcelona)
 793 meteorological and solar radiation sites are closely located.

794

795 Figure 3. Spatial variability of mean June–August O_3 (a) and O_x (b) concentrations from 12:00 to 19:00 h
 796 observed in selected air quality monitoring sites. Data from Ciutadella (CTL), Palau Reial (PLR), Montcada
 797 (MON), Granollers (GRA), Montseny (MSY), Tona (TON), Vic (VIC), Manlleu (MAN), Pardines (PAR), Montsec
 798 (MSC), Begur (BEG), Bellver de Cerdanya (BdC), Berga (BER), Agullana (AGU), Santa Pau (STP), Mataró (MAT),
 799 Manresa (MNR), Ponts (PON), Sort (SOR), Juneda (JUN), La Sénia (LSE), Constantí (CON), Gandesa (GAN),
 800 Vilanova i la Geltrú (VGe) and Alcover (ALC) air quality monitoring stations.

801

802 Figure 4. (a) Annual average traffic entering Barcelona City during weekdays (weekends not considered) during
 803 2005–2016 versus GHG emissions (attributed to industry and power generation sectors) in Catalonia during
 804 2005–2016. (b) Annual NO_x measured at CTL (Ciutadella) and MON (Montcada) air quality monitoring sites
 805 versus annual OMI-NASA’s measured background NO_2 during 2005–2017.

806

807 Figure 5. Monthly hourly average concentrations of O_3 (a) and O_x (b) along the S–N axis during 2005–2017.
 808 Data from Ciutadella (CTL), Montcada (MON), Granollers (GRA), Montseny (MSY), Tona (TON), Vic (VIC),
 809 Manlleu (MAN) and Pardines (PAR) air quality monitoring stations.

810

811 Figure 6. Monthly weekday average concentrations of O_3 concentrations calculated between 12:00 and 19:00
 812 h along the S–N axis during 2005–2017. Data from Ciutadella (CTL), Montcada (MON), Granollers (GRA),
 813 Montseny (MSY), Tona (TON), Vic (VIC), Manlleu (MAN) and Pardines (PAR) air quality monitoring stations.

814

815 Figure 7. (a) July O_3 and (b) O_x daily cycles plotted from mean hourly concentrations measured in air quality
 816 monitoring sites located along the S–N axis during 2005–2017. The black arrows point to the O_3 and O_x maxima
 817 time of the day. Data from Ciutadella (CTL), Montcada (MON), Granollers (GRA), Montseny (MSY), Tona (TON),
 818 Vic (VIC), Manlleu (MAN) and Pardines (PAR) air quality monitoring stations.

819

820 Figure 8. For the period 2005–2017, trends of the EHITs measured by air quality monitoring stations along the
 821 S–N axis (a) Annual trends of the EHITs, average temperatures measured in Vic (Gurb) (July during 13:00 to
 822 16:00 h), background NO_2 measured by OMI-NASA (June to August) and average solar radiation measured at
 823 Girona and Barcelona (June to August). (b) Monthly patterns of the EHITs, average temperatures measured in
 824 Vic, background NO_2 measured by OMI and solar radiation measured at Girona and Barcelona. (c) Weekly
 825 patterns of the EHITs and background NO_2 measured by OMI. (d) Hourly patterns of the EHITs. Despite the
 826 incomplete data availability in MAN 2005, almost 20 EHITs were recorded. AQ data from Ciutadella (CTL),
 827 Montcada (MON), Granollers (GRA), Montseny (MSY), Tona (TON), Vic (VIC), Manlleu (MAN) and Pardines
 828 (PAR) monitoring stations.



829
 830 Figure 9. Average hourly O_3 concentrations for all days with EHIT records and those without for Tona (TON),
 831 Vic (VIC), Manlleu (MAN) and Pardines (PAR) air quality monitoring stations, (left top) as well as for the NO_2
 832 levels at TON (left bottom). Average hourly increments of O_3 concentrations for all days with and without EHIT
 833 records (right); in all cases for June–August 2005–2017.

834
 835 Figure 10. Idealized two-dimensional section of O_3 circulations in the coastal region of Barcelona to the Pre-
 836 Pyrenees on a typical summer day (upper) and night (bottom). The gray shaded shape represents a
 837 topographic profile south to north direction, from the Mediterranean Sea to the south slopes of the Pre-
 838 Pyrenean Ranges (i.e., along the S–N axis). The colored dots and abbreviations depict the air quality monitoring
 839 stations located along the S–N axis: Ciutadella (CTL), Montcada (MON), Granollers (GRA), Montseny (MSY),
 840 Tona (TON), Vic (VIC), Manlleu (MAN) and Pardines (PAR). Modified and adapted to the S–N axis from Millán
 841 et al. (1997, 2000), Querol et al. (2017, 2018).

842
 843 Figure 11. Daily average background NO_2 levels in Western Europe (top) and Catalonia (bottom), June to
 844 August 2005–2017 in two different scenarios. (Left) P25: days when the maximum daily 8-h mean O_3
 845 concentrations in the Vic Plain are below the percentile 25 ($<105 \mu g m^{-3}$) and (right) P75: same but
 846 concentrations being above the percentile 75 ($>139.5 \mu g m^{-3}$).

847
 848 Figure 12. Box plots of O_x measured in TON per weekday June and July 2005–2017 for days with $\delta O_{x, TON-CTL} > 0$
 849 ($n = 793$ days with valid data). Each box represents the central half of the data between the lower quartile
 850 (P25) and the upper quartile (P75). The line across the box displays the median value. The whiskers that extend
 851 from the bottom and the top of the box represent the extent of the main body of data. The outliers are
 852 represented by black points.

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FIGURES

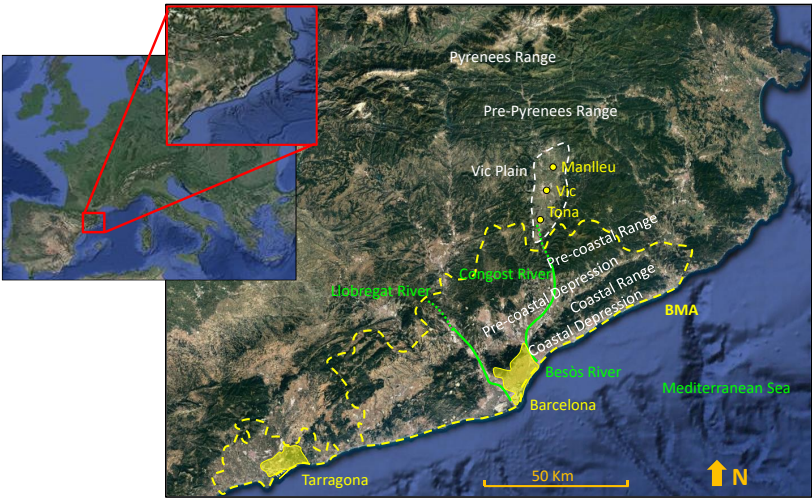


FIGURE 1

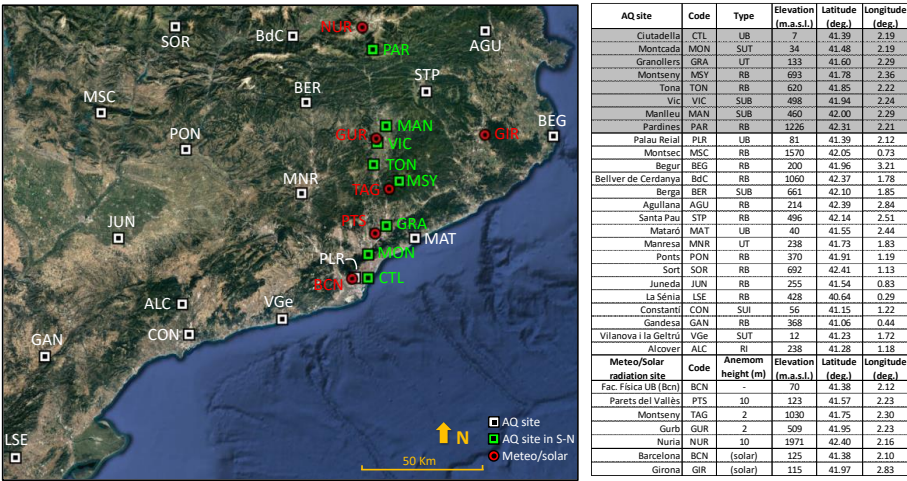
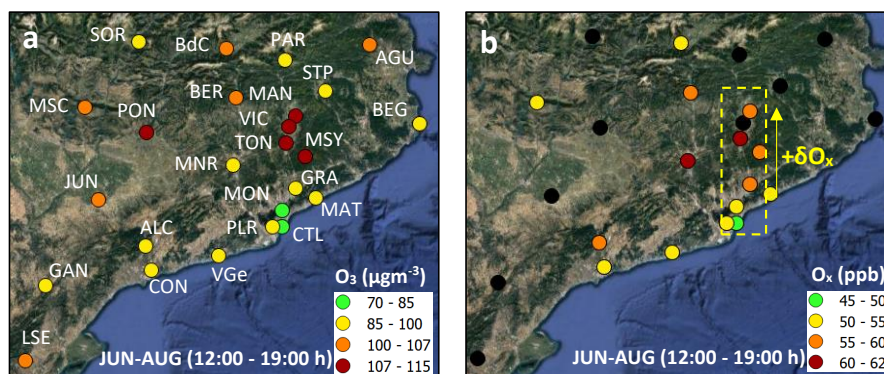
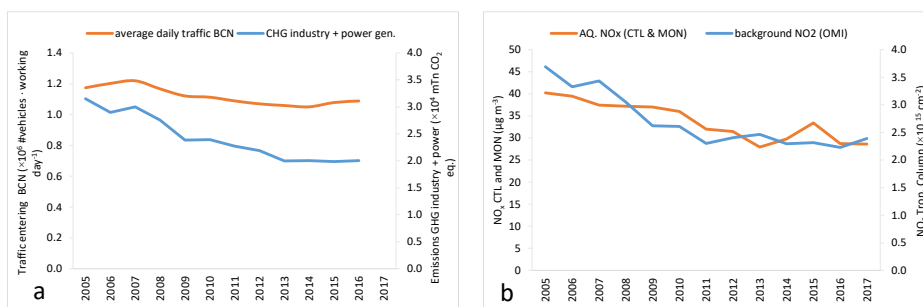


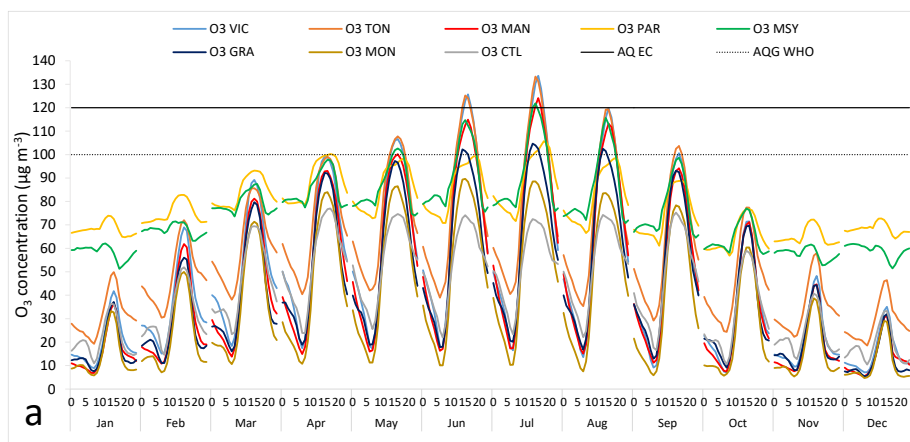
FIGURE 2



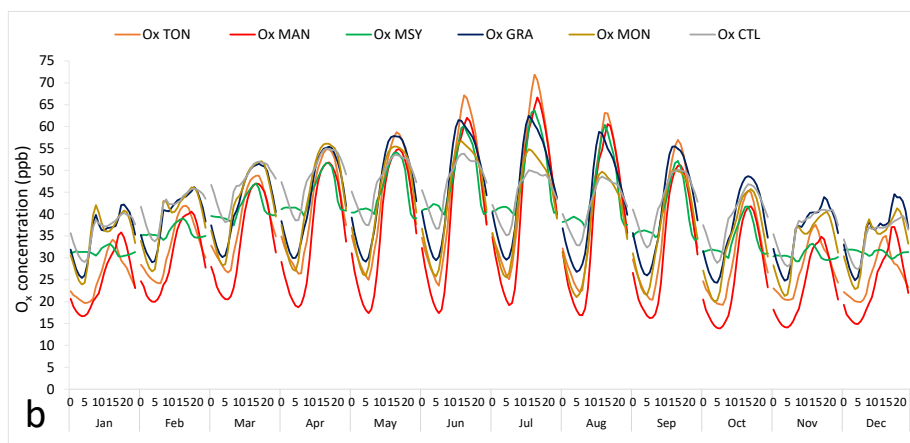
886 FIGURE 3



887 FIGURE 4

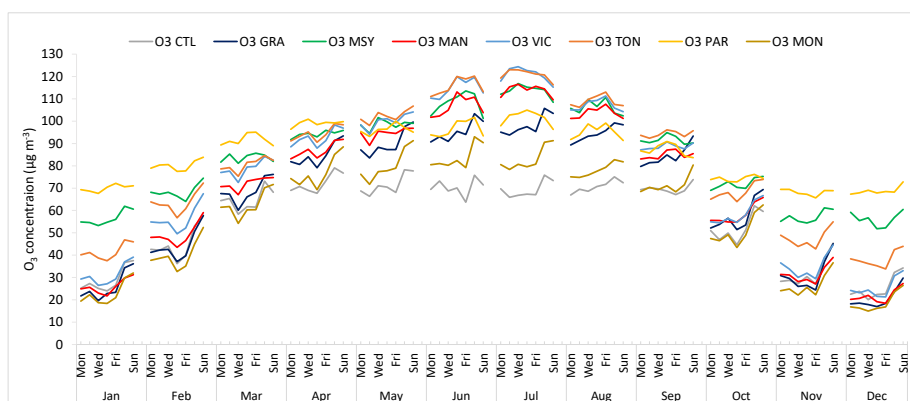


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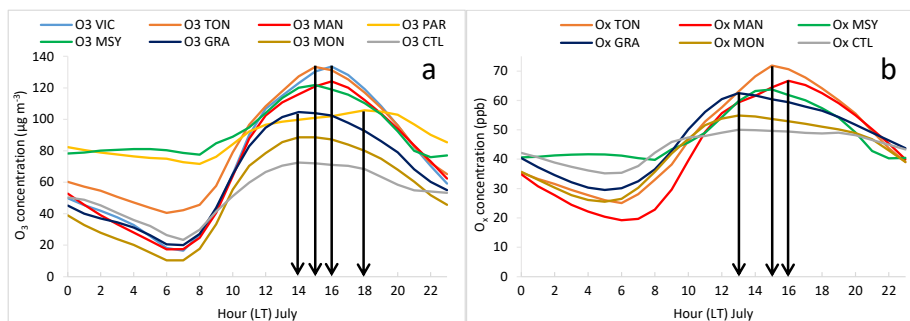
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890 FIGURE 5

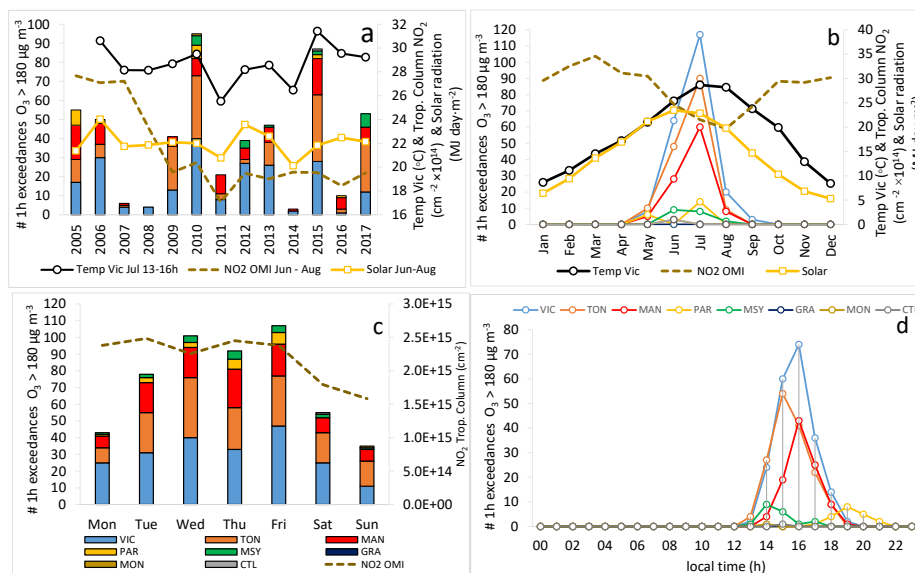


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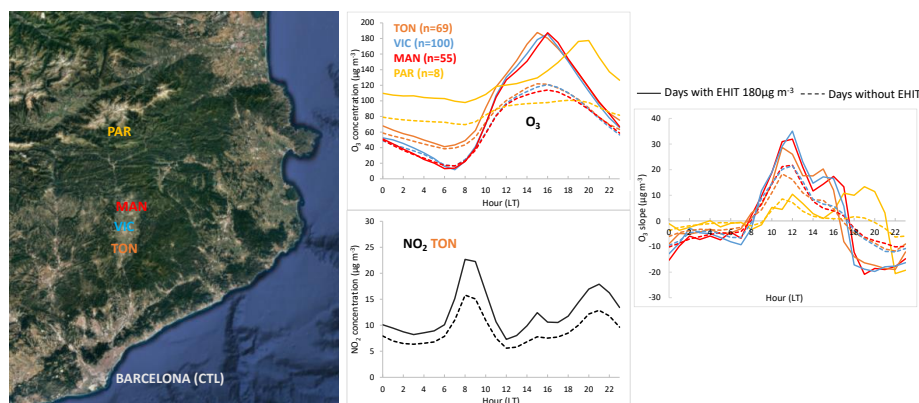
892 FIGURE 6



893 FIGURE 7



894 FIGURE 8



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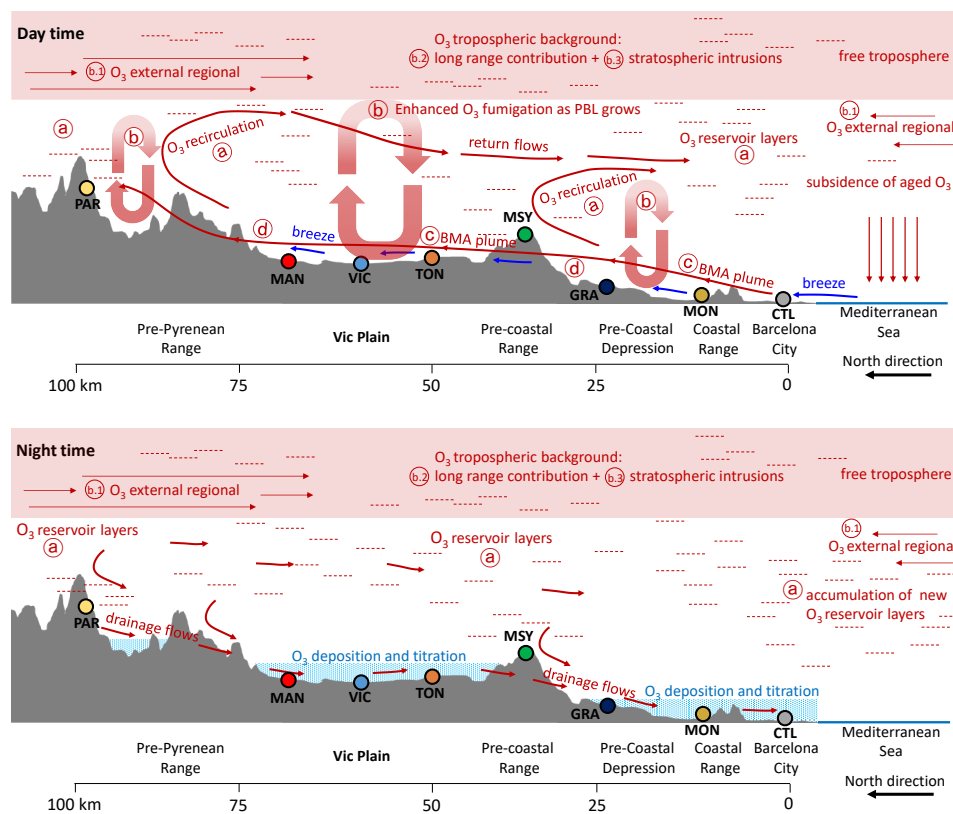


FIGURE 10

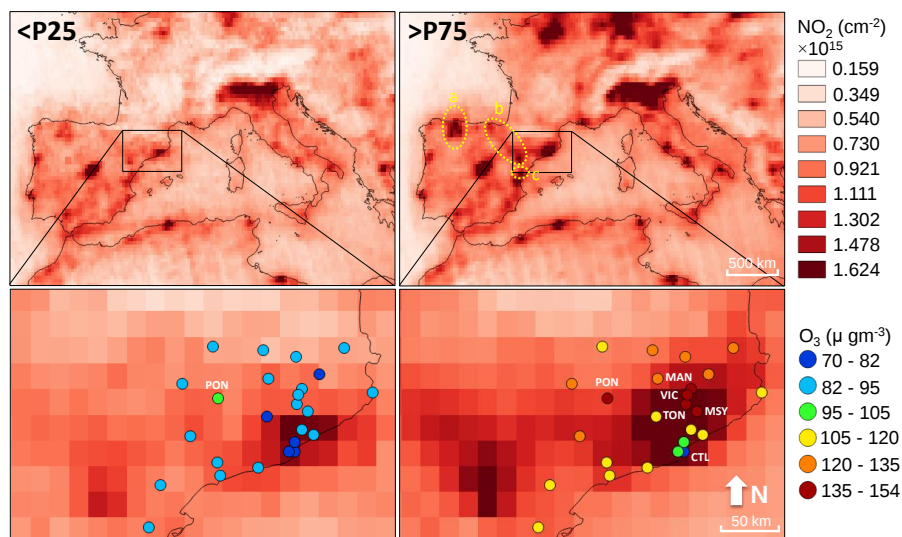
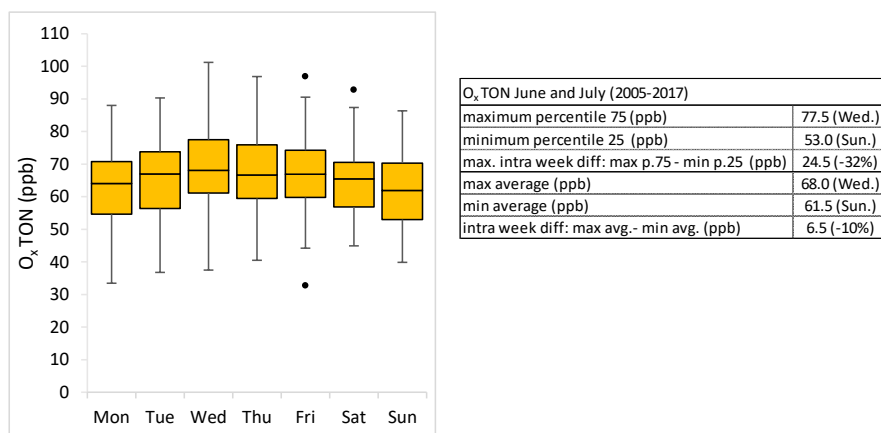


FIGURE 11



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909 FIGURE 12

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TABLE

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912 Table 1. Results of the time trend assessment carried out for annual season averages (April–September) of
 913 NO, NO₂, O₃ and O_x levels using the Theil–Sen statistical estimator. The symbols shown for the p-values related
 914 to how statistically significant the trend estimate is: p < 0.001 = *** (highest statistical significance), p < 0.01
 915 = ** (mid), p < 0.05 = * (moderate), p < 0.1 = + (low). No symbol means lack of significant trend. Units are µg
 916 m⁻³. Shaded air quality monitoring sites belong to the S–N axis. Types of air quality monitoring sites are urban
 917 (traffic or background: UT, UB), suburban (traffic, industrial or background: SUT, SUI, SUB) and rural
 918 (background: RB).

NO APR-SEP	type	period	years	initial concentration	%/year	%/year min	%/year max	%/period	units/year	units/year min	units/year max	units/period	p-Value
CTL	UB	2005-2017	13	9.7	-2.2	-3.7	-1.1	-28.8	-0.2	-0.4	-0.1	-2.9	*
MON	SUT	2005-2017	13	20.3	-4.3	-6.2	-3.3	-56.5	-0.8	-1.3	-0.5	-10.1	***
MAN	SUB	2006-2017	12	3.0	-1.1	-3.6	3.3	-13.3	0.0	-0.1	0.1	-0.3	
TON	RB	2005-2017	13	2.9	-2.9	-5.3	1.6	-37.1	-0.1	-0.1	0.0	-0.8	
BdC	RB	2005-2017	13	1.6	1.5	-1.6	11.3	20.0	0.0	0.0	0.1	0.3	
BER	SUB	2008-2017	10	4.5	1.5	-6.2	36.0	15.3	0.1	-0.4	0.5	0.6	
MAT	UB	2006-2017	12	3.0	-0.1	-2.2	2.7	-1.5	0.0	-0.1	0.1	0.0	
MNR	UT	2005-2017	13	11.8	-3.8	-5.9	-1.6	-49.0	-0.3	-0.7	-0.1	-4.5	**
CON	SUI	2005-2017	13	3.5	-2.8	-6.7	4.0	-36.7	-0.1	-0.2	0.1	-1.0	
VGe	SUT	2005-2017	13	4.8	-4.4	-6.5	-0.4	-57.2	-0.2	-0.3	0.0	-2.0	*
ALC	SUI	2005-2017	13	2.8	3.3	-1.3	21.3	42.9	0.1	0.0	0.2	1.0	
NO ₂ APR-SEP	type	period	years	initial concentration	%/year	%/year min	%/year max	%/period	units/year	units/year min	units/year max	units/period	p-Value
CTL	UB	2005-2017	13	42.1	-1.6	-2.8	-0.9	-20.6	-0.7	-1.3	-0.4	-8.5	***
MON	SUT	2005-2017	13	36.2	-2.0	-3.0	-1.2	-25.5	-0.7	-1.2	-0.4	-9.7	**
MAN	SUB	2006-2017	12	15.7	-1.3	-2.2	-0.4	-15.8	-0.2	-0.4	-0.1	-2.4	*
TON	RB	2005-2017	13	12.4	-1.4	-6.1	3.0	-18.4	-0.1	-0.7	0.2	-1.7	
BER	SUB	2008-2017	10	16.7	-2.6	-7.1	7.9	-25.6	-0.4	-1.3	0.8	-4.2	
MAT	UB	2006-2017	12	22.2	-1.7	-2.5	-0.4	-19.9	-0.4	-0.6	-0.1	-4.4	*
MNR	UT	2005-2017	13	31.0	-1.4	-2.3	-0.4	-18.0	-0.4	-0.7	-0.1	-5.4	**
CON	SUI	2005-2017	13	24.5	-1.4	-6.2	1.2	-17.6	-0.2	-1.2	0.2	-2.8	
VGe	SUT	2005-2017	13	19.0	-3.3	-4.3	-2.5	-43.3	-0.7	-0.9	-0.5	-8.6	**
ALC	SUI	2005-2017	13	10.3	-1.6	-2.5	0.0	-21.4	-0.2	-0.4	0.0	-2.6	+
O ₃ APR-SEP	type	period	years	initial concentration	%/year	%/year min	%/year max	%/period	units/year	units/year min	units/year max	units/period	p-Value
CTL	UB	2005-2017	13	45.1	1.4	-0.6	2.4	17.8	0.6	-0.3	1.1	8.3	+
MON	SUT	2005-2017	13	41.5	3.2	1.3	6.0	42.1	1.3	0.6	2.2	16.6	**
MAN	SUB	2006-2017	12	58.0	1.3	0.0	2.5	15.9	0.7	0.0	1.3	8.9	*
TON	RB	2005-2017	13	73.0	0.3	-0.5	1.4	3.7	0.2	-0.4	0.9	2.7	
VIC	SUB	2005-2017	13	61.7	0.3	-0.3	1.7	4.1	0.2	-0.2	1.0	2.5	
BEG	RB	2005-2017	13	88.5	-0.5	-1.0	0.0	-6.1	-0.4	-0.9	0.0	-5.5	
BdC	RB	2005-2017	13	76.7	-1.6	-2.1	0.0	-20.5	-1.2	-1.8	0.0	-15.9	+
BER	SUB	2008-2017	10	72.3	0.7	-0.2	2.5	7.2	0.5	-0.2	1.8	5.4	
AGU	RB	2005-2017	13	88.9	-1.1	-1.6	-0.2	-13.8	-0.9	-1.4	-0.1	-12.2	*
STP	RB	2005-2017	13	68.7	-1.4	-2.6	-0.8	-18.0	-0.9	-1.8	-0.5	-12.3	***
MAT	UB	2006-2017	12	71.6	0.4	-0.3	1.3	4.9	0.3	-0.2	0.9	3.5	+
MNR	UT	2005-2017	13	45.6	2.6	1.8	3.5	33.7	1.3	0.9	1.6	16.3	***
PON	RB	2005-2017	13	72.0	-0.1	-0.8	1.1	-1.5	-0.1	-0.6	0.8	-1.1	
LSE	RB	2005-2017	13	92.8	-0.2	-0.7	0.3	-2.7	-0.2	-0.7	0.3	-2.5	
CON	SUI	2005-2017	13	58.9	0.2	-0.4	1.6	2.9	0.2	-0.2	1.0	2.0	
GAN	RB	2005-2017	13	74.3	0.5	-0.4	2.0	6.1	0.4	-0.4	1.6	5.1	
VGe	SUT	2005-2017	13	65.6	0.4	0.0	0.9	5.4	0.3	0.0	0.6	3.6	
ALC	SUI	2005-2017	13	73.9	0.5	0.0	1.5	6.8	0.4	0.0	1.1	5.1	*
O _x APR-SEP	type	period	years	initial concentration	%/year	%/year min	%/year max	%/period	units/year	units/year min	units/year max	units/period	p-Value
CTL	UB	2005-2017	13	45.8	-0.3	-1.4	0.6	-3.6	-0.1	-0.7	0.3	-1.7	
MON	SUT	2005-2017	13	39.5	0.5	-0.6	1.7	7.0	0.2	-0.3	0.6	2.7	
MAN	SUB	2006-2017	12	37.3	0.5	-0.3	1.6	5.4	0.2	-0.1	0.5	2.0	
TON	RB	2005-2017	13	43.2	-0.2	-1.2	1.1	-2.0	-0.1	-0.6	0.4	-0.9	
BER	SUB	2008-2017	10	45.0	0.4	-0.9	1.4	4.3	0.2	-0.4	0.6	1.9	
MAT	UB	2006-2017	12	47.2	0.0	-0.5	1.1	0.5	0.0	-0.2	0.5	0.2	
MNR	UT	2005-2017	13	39.1	1.1	0.6	1.5	14.2	0.4	0.2	0.6	5.5	***
CON	SUI	2005-2017	13	42.5	-0.1	-0.5	0.7	-1.5	0.0	-0.2	0.3	-0.6	
VGe	SUT	2005-2017	13	43.1	-0.5	-0.9	0.1	-6.5	-0.2	-0.4	0.0	-2.9	+
ALC	SUI	2005-2017	13	42.5	-0.1	-0.4	1.0	-1.0	0.0	-0.2	0.4	-0.5	

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