



1 **Is the photochemistry activity weak during haze events?**
2 **—— A novel exploration on the photoinduced heterogeneous reaction of**
3 **NO₂ on mineral dust**

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10 **Abstract**

11 Despite the increased awareness of heterogeneous reaction on mineral dust, the knowledge of how the
12 intensity of solar irradiation influences the photochemistry activity remains a crucially important part in
13 atmospheric research. Relevant studies have not seriously discussed the photochemistry under weak sunlight
14 during haze, and thus ignored some underlying pollution and toxicity. Here, we investigated the heterogeneous
15 formation of nitrate and nitrite under various illumination conditions by laboratory experiments and field
16 observations. Observed by *in-situ* diffuse reflectance infrared Fourier transform spectroscopy (DRIFTS),
17 water-solvated nitrate was the main surface product, followed by other species varying with illumination condition.
18 The growth of nitrate formation rate tends to be slow after the initial fast with increasing light intensity. For
19 example, the geometric uptake coefficient (γ_{geo}) under 30.5 mW/cm² (5.72×10^{-6}) has exceeded the 50 % of that
20 under 160 mW/cm² (1.13×10^{-5}). This case can be explained by the excess NO₂ adsorption under weak illumination
21 while the excess photoinduced active species under strong irradiation. Being negatively associated with nitrate
22 ($R^2=0.748$, $P<0.01$), nitrite acts as the intermediate and decreases with increasing light intensity via oxidation
23 pathways. Similar negative dependence appears in coarse particles collected during daytime ($R^2=0.834$, $P<0.05$),
24 accompanied by the positive association during nighttime ($R^2=0.632$, $P<0.05$), suggesting illumination a
25 substantial role in atmospheric nitrogen cycling. Overall, for the nitrate formation, the conspicuous response under
26 slight illumination offers opportunities to explain the secondary aerosol burst during haze episodes with weak
27 irradiation. Additionally, high nitrite levels accompanied by low nitrate concentrations may induce great health
28 risk which was previously neglected. Further, Monte Carlo simulation coupled with sensitivity analysis may
29 provide a new insight in the estimations of kinetics parameters for atmospheric modelling studies.



30 1 Introduction

31 Secondary nitrate aerosols, deriving mainly from various oxidation processes of nitrogen oxides, are of great
32 importance in atmospheric chemistry (Anenberg et al., 2017). These ubiquitous species is key for describing the
33 composition and sources of particulate matters (Huang et al., 2014; Schuttlefield et al., 2008). It was
34 investigated that the contributions of nitrate to the particle mass concentration increase throughout the pollution
35 episodes (Guo et al., 2014). However, current atmospheric models fail to capture the serve nitrate enhancement
36 from the clean to haze period, and thus triggers the discussion on the heterogeneous reaction of NO₂ on primary
37 aerosols (Tan et al., 2016). Modeling studies indicated that nitrate formation is highly associated with airborne
38 mineral dust (Tan et al., 2016). Accounting for ~36% of the total primary aerosol emissions, mineral dust is one
39 of the most abundant particle types in the troposphere (Chen et al., 2012; Usher et al., 2003). During their global
40 journey, many heterogeneous reactions take place on the particle surface, and further affect the atmospheric
41 oxidation capacity (Tang et al., 2017). This process has aroused widespread interest in East Asia because dust
42 occupies a great share in fine particles due to the frequent occurrence of sand storms (Zhang et al., 2015). Hence,
43 the heterogeneous reaction of NO₂ on mineral dust is worthy of broader concerns.

44 Titanium dioxide (TiO₂) is found in mineral dust at mass mixing ratios ranging from 0.1% to 10% depending
45 on the exact location from where the particles were uplifted (Ndour et al., 2008). Compared to other
46 non-semiconducting components in mineral dusts, TiO₂ has direct environmental implications for its
47 photocatalysis (Nanayakkara et al., 2014). Prior studies have indicated the photoinduced oxidation of trace gases
48 by TiO₂ an essential role in the chemical balance of the atmosphere (Chen et al., 2012). Additionally,
49 TiO₂-coating surfaces are currently used on building exteriors, road lamps and road bricks (Ballari et al., 2010;
50 Ballari et al., 2011). These self-cleaning materials in populated urban areas facilitate the irreversible removal of
51 NO₂ from atmosphere with the substantial formation of gaseous nitrous acid (HONO) and ozone (Balajka et al.,
52 2018; Langridge et al., 2009; Monge et al., 2010). Accordingly, TiO₂ is frequently adopted as the reference
53 material on behalf of the ubiquitous semiconducting components in atmospheric environment, especially the urban
54 atmosphere.

55 For the heterogeneous process on mineral dust, prior studies put close attention to varied influential factors.
56 Among these, moisture and temperature are widely concerned and significant advances have been made (Li et
57 al., 2010; Tan et al., 2017; Tan et al., 2016; Wang et al., 2012). Although being treated as an important index in
58 many atmospheric discussions, illumination has not been systematically investigated for its effects on the
59 heterogeneous uptake of trace gases. Most remarkable studies concerned the photocatalytic effects instead of the



60 dependence on illumination conditions (Dupart et al., 2014; Guan et al., 2014). Some researchers (El Zein and
61 Bedjanian, 2012) measured the reactive uptake coefficients (γ -values) for the heterogeneous reaction of NO₂ on
62 TiO₂ under various irradiance intensities while ignored the reaction mechanism behind the variation. Furthermore,
63 nitrite is of great significance in atmospheric processes for its frequent appearance and great contributions to
64 aerosol toxicity. However, there is little information available in literature about the pollution characteristics or
65 reaction pathways of nitrite aerosols. Generally, how the illumination influences the uptake capacity and product
66 species are problems urgently needs solving.

67 This work aims to provide a fresh perspective to explore the light dependence for the heterogeneous reaction
68 on mineral dust. Monte Carlo simulation is introduced to evaluate the kinetics for nitrate formation. Atmospheric
69 particulates were collected and analyzed to support relevant findings. This research could help further understand
70 the illumination effects in the atmospheric nitrogen cycling, and simultaneously provide extremely valid
71 parameters for modelling studies.

72 2 Experimental

73 2.1 Materials

74 Commercial TiO₂ (Degussa, Germany), with an anatase-to-rutile ratio of 3: 1, an average particle size of
75 12.05 ± 3.46 nm and a Brunauer-Emmett-Teller (BET) specific surface area (S_{BET}) of 55.83 ± 0.35 m²·g⁻¹ was
76 employed as the photocatalytic mineral dust (SI, Section SI). All chemicals were of analytical grade and obtained
77 from Aladdin Chemical Reagent Co., Ltd. Water in all experiments was ultrapure water (specific resistance ≥ 18.2
78 M Ω cm).

79 High-pure air (79% N₂ and 21% O₂, Shanghai TOMOE Co., LTD, China) and 300 parts per million (ppm)
80 NO₂ (N₂ dilution, Shanghai Qingkuan Co., LTD, China) were included in this research. Prior to coming into the
81 gas supply system, high-pure air went through silica gel and molecular sieve for drying and purification.

82 2.2 DRIFTS experiments

83 A FTIR spectrometer (Tracer-100, Shimadzu, Japan) equipped with a liquid-nitrogen-cooled
84 mercury-cadmium-telluride (MCT) detector was applied to record in situ DRIFTS spectra with 100 scans averaged
85 for each spectrum and a resolution of 4 cm⁻¹. We have described the general features of the setup in Figure S2 and
86 previous reports (Wang et al., 2018a; Wang et al., 2018b; Wang et al., 2018c). Herein, a xenon lamp
87 (CEL-TCX250, Beijing Ceaulight Co., LTD, China) was used to provide simulated solar irradiation upon the
88 particles (Figure S3).



89 Prior to each experiment, the particles were pretreated in a stream of high-pure air (200 ml·min⁻¹) for 60 min
90 to remove the adsorbed water and impurities from the surfaces (**Figure S4**). Due to the overlapping bands of
91 adsorbed water (~1640 cm⁻¹) and nitrogen compounds, the sample after pretreatment was exposed to humid
92 high-pure air (RH≈30%, 100 ml·min⁻¹) for 20 min, after which the moisture absorption reaches saturation (**Figure**
93 **S5**). A background spectrum was recorded after the process and then NO₂ calibration gas (5.12 ml·min⁻¹) was
94 added into the DRIFTS chamber with a calculated concentration of 15.33 ppm. Calibration gases with NO₂
95 concentrations of 9.20 and 21.45 ppm were also involved for the concentration dependence experiments. Ten
96 light intensity levels (0.0, 0.3, 5.4, 17.5, 23.8, 30.5, 54.5, 98.5, 128.1, and 160.0 mW·cm⁻²) were referred in this
97 study.

98 Each test lasted 90 min, during which a series of spectra were recorded every 5 min. The reacted particles
99 were extracted by oscillation (5 min) with 4 ml water. The extraction solution was then passed through a 0.22 μm
100 PTFE membrane filter for ion detection.

101 2.3 Ion analysis

102 The nitrate and nitrite ions were analyzed by an ion chromatography (IC, 883 Basic, Metrohm, Switzerland),
103 which consists of an analytical column (A5-250) and a guard column. The detection was conducted by using 3.2
104 mmol·L⁻¹ Na₂CO₃ and 1.0 mmol·L⁻¹ NaHCO₃ at a stable flow rate of 0.70 ml·min⁻¹. Multipoint calibrations were
105 performed by means of standard solutions. Good linearity of the calibration curve was obtained with $R^2 > 0.998$.

106 2.4 Photo-electrochemical (PEC) test

107 In order to qualitatively evaluate the generation of electron-hole pairs under different light intensities, PEC
108 tests were conducted by a electrochemical workstation (CHI-660D, Shanghai Chenhua Co., LTD, China) in a
109 three-electrode cell with a quartz window (**Yang et al., 2017; Zheng et al., 2015**). TiO₂ particles were deposited
110 on a sheet of fluorine-tin-oxide glass to serve as the working electrode with an effective area of 1cm². A platinum
111 wire and an Ag/AgCl electrode were employed as the counter and reference electrodes, respectively. The
112 electrolyte was 0.5 mol/L NaNO₃. A xenon lamp (CEL-S500, Beijing Ceaulight Co., LTD, China) was used to
113 provide simulated sunlight.

114 2.5 Uptake coefficient estimation

115 The reactive uptake coefficient, γ , is defined as the ratio of the reactive gas-surface collision rate ($d[NO_3^-]/dt$)
116 to the total gas-surface collision rate (Z) (**Gustafsson et al., 2006**). The equations are shown as follows.

$$117 \quad \gamma = \frac{d[NO_3^-]/dt}{Z} \quad Eq. (1)$$



$$\frac{d[NO_3^-]}{dt} = slope \times f \quad Eq. (2)$$

$$Z = \frac{1}{4} \times A_s \times [NO_2] \times v_{NO_2} \quad Eq. (3)$$

$$v_{NO_2} = \sqrt{\frac{8RT}{\pi M_{NO_2}}} \quad Eq. (4)$$

Where *slope* represents the growth rate of the nitrate peaks, *f* is the conversion factor, *A_s* is the particle reactive surface area, *v_{NO₂}* is the mean velocity of NO₂ molecule, [NO₂] is the NO₂ concentration, *R* is the gas constant, *T* is the temperature, *M_{NO₂}* is molecular weight of NO₂ (**Table S1**).

The conversion factor (*f*) is obtained from a calibration plot with the amount of NO₃⁻ versus the integrated areas for nitrate (**Tan et al., 2017; Tan et al., 2016**). The factor is 2.09×10¹⁵±1.61×10¹⁴ (ion · K-M unit⁻¹) in this study (**Figure S6**). For the *A_s*, both geometric surface area (*A_{geo}*) and BET surface area (*A_{BET}*) are mentioned to evaluate the upper and lower limits of the *γ*-values (denoted as *γ_{geo}* and *γ_{BET}*, respectively) varying with reaction probabilities between reactants and particles.

Monte Carlo simulation was implemented to deal with the uncertainties (**Chiang et al., 2009; Ginsberg and Belleggia, 2017; Xia et al., 2013**). Each independent variable was determined via five or more replication measurements and assumed to be normally distributed in the simulation. Based on earlier finds, 5000 iterations are sufficient to ensure the stability of the results. Additionally, sensitivity analysis is helpful in exploring the variables that influence the estimation most. Pearson correlation coefficients between each variable and the output (*γ*-value) were calculated and then normalized to 100%. On this basis, the contribution of each input variable to the output can be assessed. Three input variables are included for *γ_{geo}*: *slope*, *f*, and *A_s*. For *γ_{BET}*, the *A_s* is further divided into mass and *S_{BET}* as discussed above.

2.6 Particle sampling and chemical analysis

Aerosols were collected in the late summer and early autumn in the campus of Fudan University, Shanghai, China (**Figure S8**). The first stage from 23th August to 17th September contains 26 daily samples. The second stage lasted from 21th to 29th September, including eight sample sets collected during daytime and another eight during nighttime.

The size-segregated samples ranging from 0.4 to 100 μm were collected on quartz fiber filters (Whatman, UK) using an eight stage micro-orifice uniform deposit impactor (Anderson, Tisch Environmental Inc, USA) operating at a flow rate of 28.3 L/min. The particle modes were defined as follows: 0-0.56 μm for condensation mode, 0.56-1.8 μm for droplet mode, and 1.8-100 μm for coarse mode.



146 Before sampling, the filters were pre-combusted at 550°C for 4 h to minimize original impurities. After
147 collection, the filters were extracted ultrasonically by 20 ml water for 45 min. Water extracts were passed through
148 a 0.22 µm PTFE membrane filter for NO_3^- and NO_2^- detection as introduced in **section 2.3**.

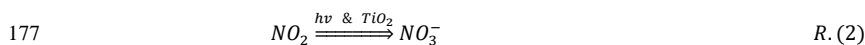
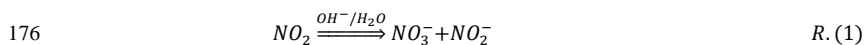
149 **3 Results and discussion**

150 **3.1 Observed species on particles**

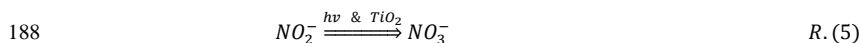
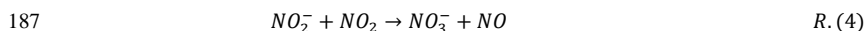
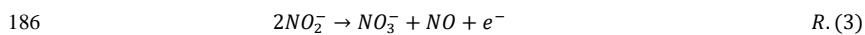
151 **Figure 1** presents the DRIFTS spectra recorded in the absence and presence of illumination, coupled with the
152 Gaussian curve-fitting procedure to deconvolute the overlapping bands. The fitting was undertaken until
153 reproducible results were obtained with the coefficient of determination (R^2) greater than 0.990. The bands in the
154 spectra are quite rich, indicating various products as summarized in **Table S3**.

155 Under illumination, the signals peaking at 1312 and 1553 cm^{-1} reflect the formation of monodentate nitrate,
156 whereas those at 1276, 1573, and 1602 cm^{-1} account for the vibration of bidentate nitrate (**Figure 1d, e**) (**Li et al.,**
157 **2010; Ma et al., 2011; Niu et al., 2017; Szanyi et al., 2007**). Bridging bidentate nitrate can be further identified
158 by the shoulder peak at 1602 cm^{-1} (**Du et al., 2019; Goodman et al., 1998; Sun et al., 2016**). Besides, the peaks
159 at 1347 and 1412 cm^{-1} are assigned to water-solvated nitrate (**Baltrusaitis et al., 2007; Guan et al., 2014; Miller**
160 **and Grassian, 1998**). Under dark condition, except the similar bands appearing under illumination (1561, 1409,
161 1323, and 1271 cm^{-1}), some nitrite products become more attractive as evident by the monodentate nitrite at 1195
162 and 1440 cm^{-1} , as well as the bidentate nitrite around 1308 cm^{-1} (**Figure 1a, b**) (**Wu et al., 2013**). Water-solvated
163 nitrate is far ahead in amount compared to other species (**Figure 1c, f**), suggesting weak links between the products
164 and particle surfaces. Hence, the surface water layers of the hygroscopic particles provides plenty active space for
165 the heterogeneous uptake of NO_2 .

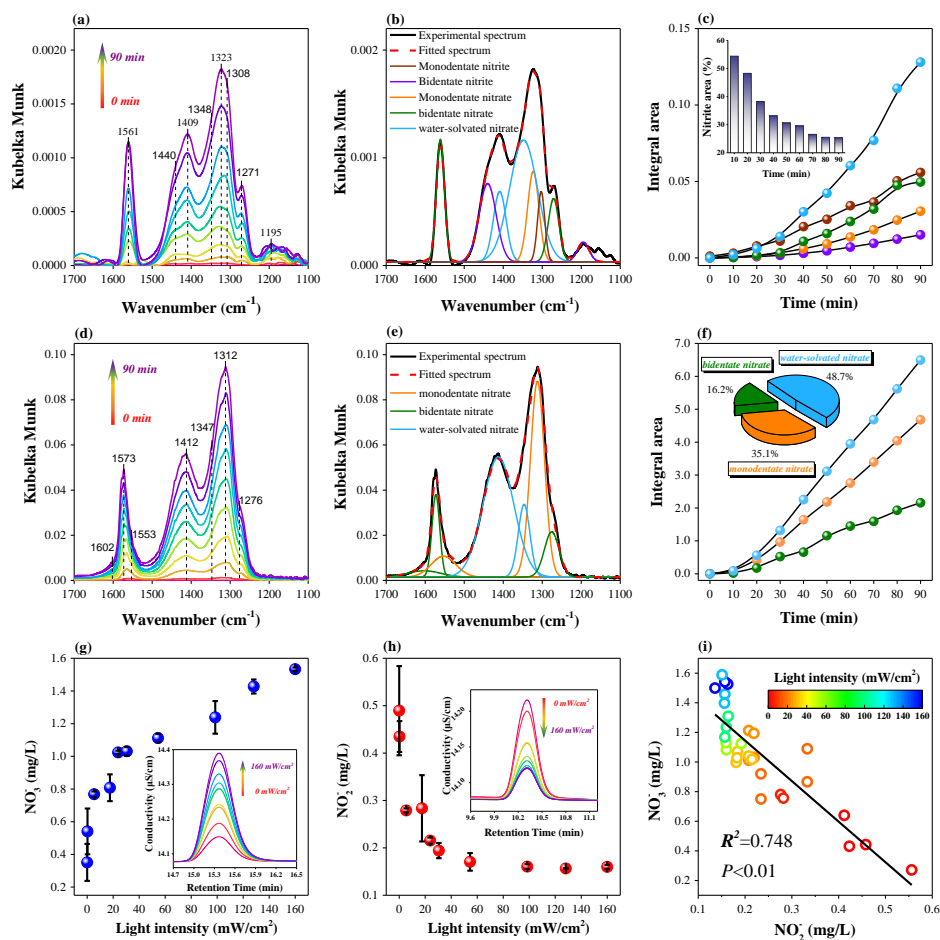
166 After the reversible adsorption of NO_2 on mineral dust (**R.S1**), the NO_2 reacts with hydroxyl-related groups
167 (OH) or surface H_2O to form adsorbed nitrate/nitrite or free nitric acid/nitrous acid, respectively (**R.S2-S3**). Since
168 no acid molecules were observed, free nitrite and nitrate ions stem from ionization (**R.S4-S5**). The
169 disproportionation process (**R.I**) dominates the dark reaction. When excited with light (wavelength \geq 390 nm),
170 there is the generation of electron-hole pairs in the conduction and valence bands of TiO_2 (**R.S6**) (**Dupart et al.,**
171 **2014; FUJISHIMA and HONDA, 1972; Yu and Jang, 2018**). Photogenerated holes and electrons react with
172 H_2O and O_2 , and thus lead to the formation of hydroxyl radicals (OH^\cdot) and reactive oxygen radicals (O_2^-),
173 respectively (**R.S7-S8**) (**Chen et al., 2012**). Superoxide hydrogen radical (HO_2^\cdot) and hydrogen peroxide (H_2O_2)
174 appear and produce OH^\cdot as well (**R.S7-S14**). These photoinduced active species (PAS) would accelerate the
175 nitrate formation (**R.2**).



178 Noticeably, nitrite (especially monodentate type) decreases in proportion as the dark reaction proceeds,
179 accompanied by the increasing contribution from bidentate nitrate species and water-solvated ones (**Figure 1c**,
180 **S10a**). The nitrite would react with another surface nitrite in a Langmuir-Hinshelwood mechanism (**R.3**) or
181 gaseous NO₂ in an Eley-Rideal mechanism (**R.4**) to form nitrate in the absence of illumination (**Tang et al., 2018**;
182 **Underwood et al., 1999**). Oxygen also acts as a promoter in the nitrite oxidation (**Tang et al., 2018**). On the other
183 hand, diverse nitrate species make steady contributions to the total products during the photoreactions (**Figure**
184 **S10b**). Generally, nitrite signal is visible in dark, while gradually fades away after irradiation due to the oxidation
185 of nitrite to nitrate by PAS via **R.5** (**Section S8**).



189 Illumination has impacts on either product species or the production. The final DRIFTS spectra grow in
190 intensity as the illumination becomes stronger. Raman measurements also indicate the drastic enhancement caused
191 by sunlight, evident by the higher nitrate peak after illumination compared to that after dark process (**Section S9**)
192 (**Fu et al., 2017**; **Yu et al., 2018**; **Zhao et al., 2018**). These observations provide a solid evidence that the nitrate
193 formation on mineral dust is enhanced under sunlight, in nice agreement with previous results (**Dupart et al.,**
194 **2014**; **Guan et al., 2014**). Noticeably, the nitrate determined by IC exhibits a clear nonlinear uptrend with
195 increasing light intensity, suggesting uneven illumination effect on nitrate formation (**Figure 1g**). On the contrary,
196 the nitrite presents a nonlinear downtrend (**Figure 1h**), and thus results in the negative association with nitrate
197 (**Figure 1i**). What is the proposed mechanism behind the uneven illumination effects? Whether the photoinduced
198 negative dependence appears in atmospheric particulates? We may discuss these issues in the following sections.



199
 200 **Figure 1.** Product observations under (a-c) dark condition and (d-f) illumination ($I=98.5 \text{ mW/cm}^2$), as well as (g-i)
 201 ion analysis results. (a, d) DRIFTS spectra of nitrate and nitrate species. (b, e) Peak fitting for the final spectra
 202 based on Gaussian method. (c, f) Integral areas of diverse species as a function of reaction time. IC measurements
 203 for (g) nitrate and (h) nitrite ions after DRIFTS tests. Error bars represent 1σ . (i) Linear association between nitrate
 204 and nitrite varying with light intensity. Inset: (c) Time-dependent contributions of nitrite, (f) Contributions of
 205 diverse products after 90 min exposure. (g, h) Conductivity spectra from IC.

206 3.2 Photoinduced uptake capacity

207 To accurately evaluate the photoinduced nitrate formation, γ -values were estimated based on Monte Carlo
 208 simulation with the cumulative probability distributions depicted in **Figure 2** and the percentile values
 209 summarized in **Table S2**. γ_{BET} and γ_{geo} exhibit similar variation trends. Since the reaction is first order with respect



210 to NO₂ concentration under various light intensities (**Figure S7**), the γ -values would still be authentic for
 211 atmospheric reactions with lower NO₂ concentrations. The growth of γ -values appears to be slow after the initial
 212 fast with increasing light intensity. For instance, the γ_{geo} under 30.5 mW/cm² exceeds the half of that under 160
 213 mW/cm². To facilitate comparison, theoretical γ -values were calculated in a linear way based on the results under
 214 0 and 160 mW·cm⁻². The actual γ -values under 5.4, 17.5, 23.8, 30.5, 54.5, 98.5, and 128.1 mW·cm⁻² are 73%,
 215 135%, 189%, 158%, 148%, 103%, 39%, and 16% higher than the corresponding theoretical ones, respectively.
 216 This ‘fast-slow’ uptrend seems to be of great importance as it shows that the γ -values measured at designed
 217 irradiation intensity may not be extrapolated in a linear way to those relevant to the atmosphere. The balance
 218 between PAS formation and NO₂ adsorption is responsible for the uneven illumination effect, which will be
 219 carefully discussed in the mechanism section.

220 To distinguish the contributions of each variable to the output, sensitivity analysis is performed on the basis of
 221 the simulated data. Slope and f contribute most to the total variance of γ_{BET} and γ_{geo} , while S_{BET} and m for γ_{BET} , and
 222 A_{geo} for γ_{geo} contribute little (**Section S10**). Accordingly, slope and f values in a more accurate level are beneficial
 223 for γ -value estimation. More attention in the future needs to be devoted to the stability of DRIFTS and IC
 224 measurements.

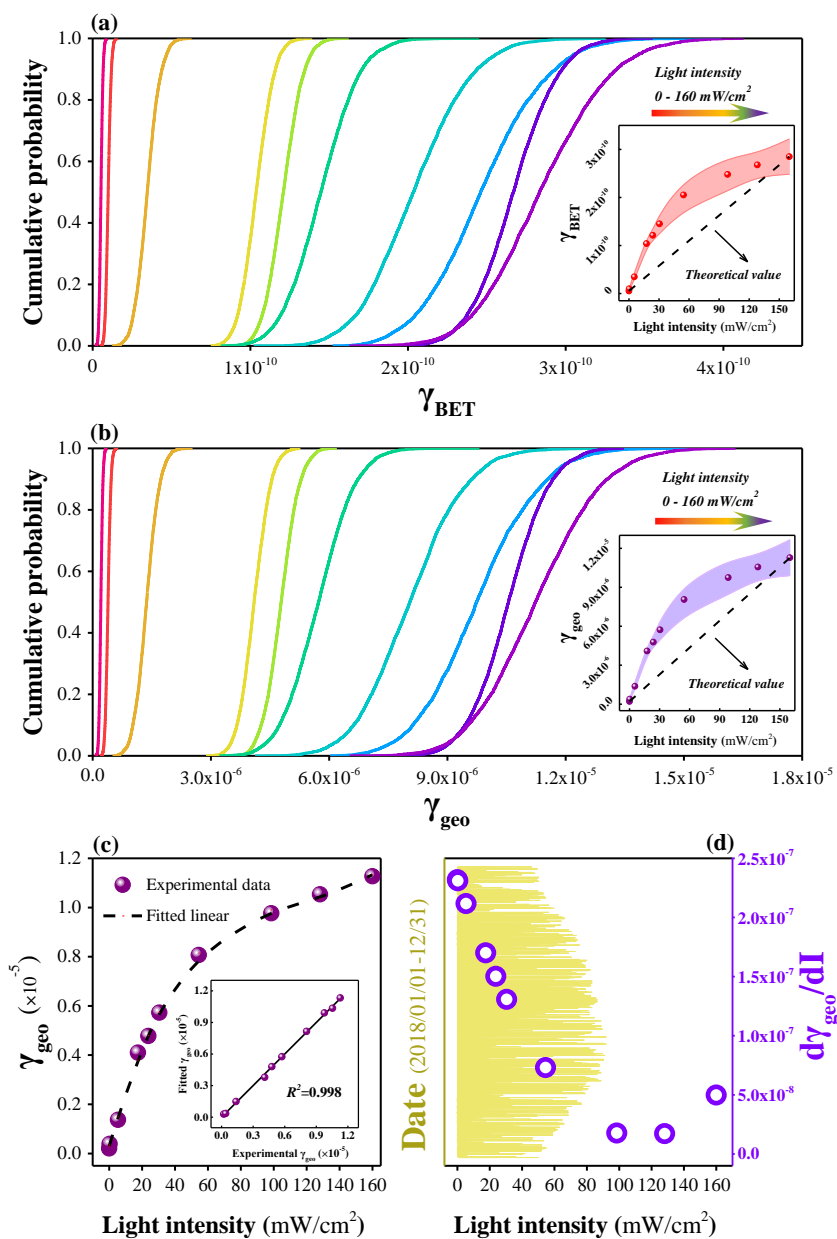
225 In view of the great significance of γ_{geo} in atmospheric models, regression analysis is employed to fit the
 226 obtained results and further predict values for relevant reactions. Since the γ -values exhibit a ‘fast-slow’ uptrend, a
 227 polynomial regression model (**Eq.5**) is used to describe the variation.

$$228 \quad \gamma_{f,\text{geo}} = aI^3 + bI^2 + cI + d \quad \text{Eq. (5)}$$

229 Where $\gamma_{f,\text{geo}}$ is the fitted γ_{geo} , I is light intensity, and a , b , c and d are essential parameters. The final formula
 230 (**Eq.6**) could explain 99.8% variation of the experimental γ_{geo} , indicating accurate regression (**Figure 2c**).
 231 Furthermore, the $d\gamma_{\text{geo}}/dI$ values are obtained by derivation to distinguish the illumination effect varying with light
 232 intensity. The uptake capacity is extremely sensitive to light under low intensity, while tends to be
 233 light-independent under strong irradiation (**Figure 2d**). In Shanghai, the 3h-average intensities are mostly lower
 234 than 80 mW/cm² (NOAA data, <https://www.arl.noaa.gov/>), indicating noticeable sunlight impacts. More
 235 importantly, the irradiation tends to be weaker in winter, highlighting the central role of light-dependent
 236 heterogeneous reaction in haze events during cold time.

$$237 \quad \gamma_{f,\text{geo}} = 5.62 \times 10^{-12} \times I^3 - 1.92 \times 10^{-9} \times I^2 + 2.32 \times 10^{-7} \times I + 2.93 \times 10^{-7} \quad \text{Eq. (6)}$$

$$238 \quad d\gamma_{\text{geo}}/dI = 1.686 \times 10^{-11} \times I^2 - 3.84 \times 10^{-9} \times I + 2.32 \times 10^{-7} \quad \text{Eq. (7)}$$



239

240

Figure 2. Cumulative probability distribution of the (a) γ_{BET} and (b) γ_{geo} values based on Monte Carlo simulation. Insets reveal the actual γ -values (Mean $\pm 1\sigma$) and theoretical ones. (c) Regression analysis on γ_{geo} .

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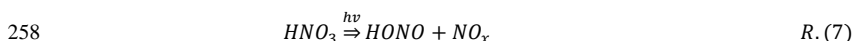
Inset presents the linear relation between calculated and fitted values. (d) Downward short wave radiation

243

flux (DSWF) in Shanghai, China coupled with estimated $d\gamma_{\text{geo}}/dI$.

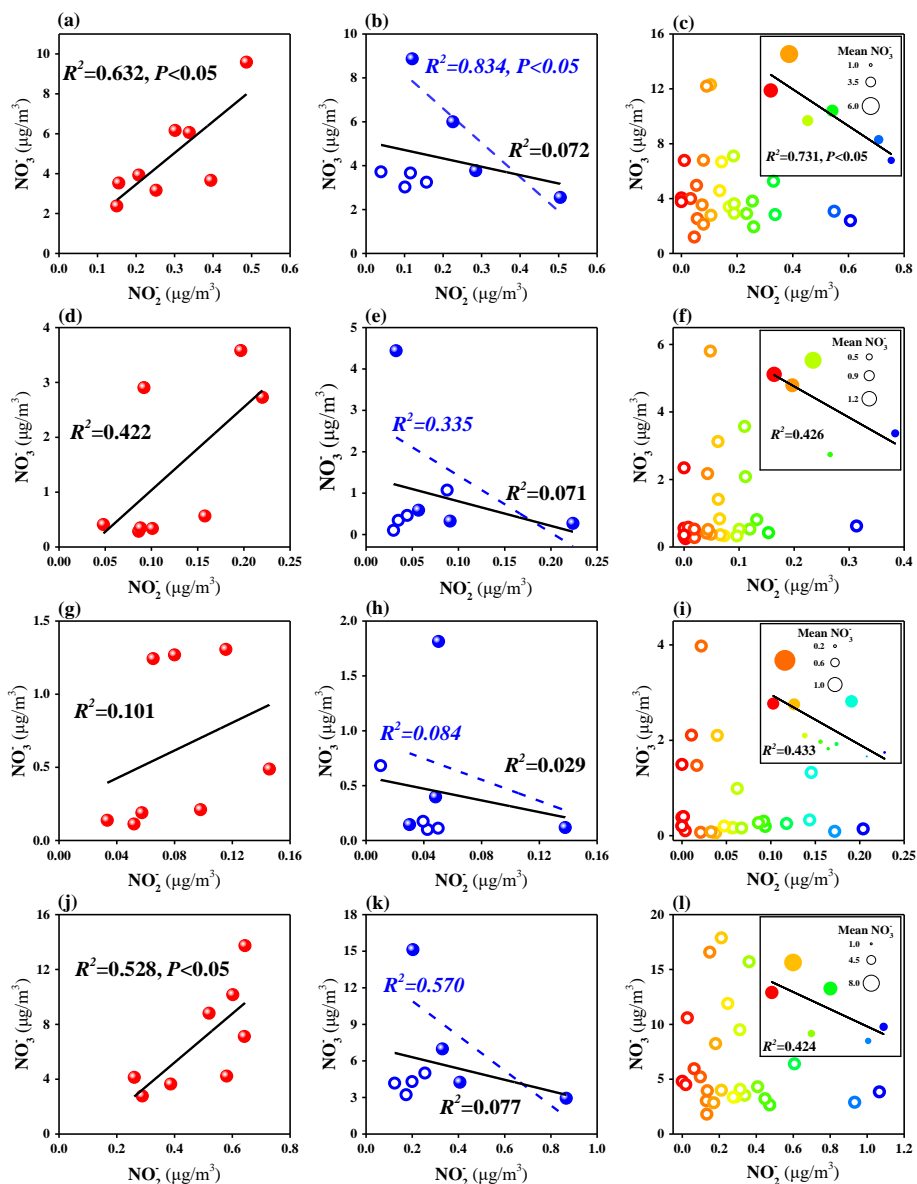
244 **3.3 Nitrogen redox**

245 **Figure 3** presents the association between atmospheric nitrate and nitrite varying with particle mode and
246 sampling period. Significant positive correlation can be found during nighttime in coarse mode (**Figure 3a**).
247 However, there is no case indicating high nitrite and nitrate levels during daytime, and the dependence seems to be
248 negative (**Figure 3b**). The correlation turns to be significant with the ignorance of cases where the nitrite and
249 nitrate concentrations are extremely low. As summarized in **section 3.1**, the associations during nighttime and
250 daytime can be explained by the NO₂ disproportionation in the absence of sunlight (**R.1**) and the nitrite oxidation
251 under illumination (**R.5**), respectively. Daily nitrite and nitrate concentrations exhibit the similar variation with
252 that during daytime, and a negative correlation can be observed based on the classification of nitrite levels (**Figure**
253 **3c**). Noticeably, low nitrite levels are usually accompanied by slight nitrate pollution in the presence of sunlight,
254 resulting mainly from the HONO formation in acidic media (**R.6**) (Liu et al., 2015; Su et al., 2011; Wang et al.,
255 2015; Zhang et al., 2012) and the photolysis of particulate nitrate (**R.7**) (Nanayakkara et al., 2014; Ostaszewski
256 et al., 2018; Schuttelfield et al., 2008; Ye et al., 2017; Ye et al., 2016).



259 Atmospheric nitrate and nitrite from diverse periods exhibit analogous size distribution: greatest in coarse
260 mode, followed by droplet mode and condensation mode (**Figure S9**). Yet, except the large mass fraction in coarse
261 mode, nitrite presents extra peak under 1.8 μm, indicating reaction pathways differing from nitrate formation
262 (Moore et al., 2004). That is, nitrate is difficult to accumulate by aqueous reactions or homogeneous processes
263 while nitrite seems to be easy, which results in the lower correlation coefficients for small size particles (**Figure**
264 **3d-i**). Since the main reaction pathways (**R.1, 3, 4**) still take place in aqueous media, and some other oxidants (e.g.
265 H₂O₂, O₃, and Fe³⁺) would replace the promoting role of semiconductor components in mineral dust under
266 illumination (**R.2, 5**) (Hems et al., 2017; Hou et al., 2017; Xue et al., 2016), the correlation in droplet mode
267 appears to be obvious with merely lower coefficients. Furthermore, both ions exhibit great mass fractions (>50%)
268 in coarse mode, making the associations for full-size particles similar with those for coarse aerosols (**Figure 3j-l**).

269 Generally, atmospheric nitrite is positively correlated with nitrate during nighttime, whereas presents
270 negative association with nitrate in the presence of irradiation. The dependence is significant in coarse mode while
271 turns to be inconspicuous in droplet mode and condensation mode.



272

273 **Figure 3.** Associations between atmospheric nitrite and nitrate ions in (a-c) coarse mode (1.8-100 μm), (d-f)274 droplet mode (0.56-1.8 μm), (g-i) condensation mode (0-0.56 μm), and (j-l) full-size particles (0-100 μm)

275 collected during (a, d, g, j) nighttime, (b, e, h, k) daytime, and (c, f, i, l) whole day. Insets: mean nitrate

276 concentrations based on nitrite classification. Linear correlation analysis (solid line) was employed for each case.

277 The dashed lines for daytime cases reveal the correlation for solid circles with hollow circles ignored.

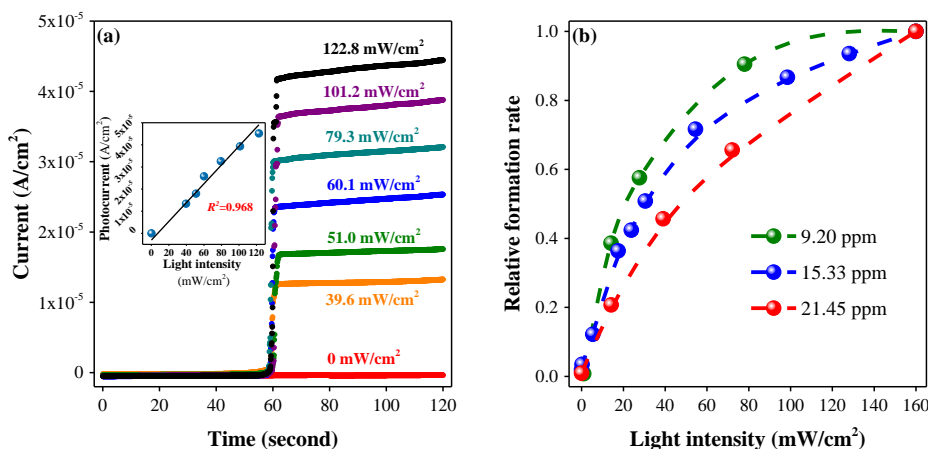


278 **3.4 Mechanism**

279 Main reaction pathways have been introduced in previous sections. However, the uneven illumination effect
280 cannot be explained by chemical equations. The photocurrent of TiO₂ is linearly correlated with light intensity
281 (**Figure 4a**), indicating even illumination effect on the production of electron-hole pairs. Hence, photocatalytic
282 activity of the mineral dust is not sufficient to explain the uneven nitrate/nitrite formation with illumination
283 variation.

284 Since no saturation effects were observed in the DRIFTS experiments, the NO₂ adsorption rate can be
285 regarded as constant. Adsorbed NO₂ becomes excess compared to the formed PAS under weak sunlight, and thus
286 makes illumination the rate-limiting factor in oxidation. At this time, nearly all the PAS participate in the oxidation
287 of surface adsorbed NO₂ as well as some nitrite intermediates. When the illumination is strong, the PAS gradually
288 become excess compared to the adsorbed NO₂. Under the circumstances, light makes little contribution to the
289 elevation of uptake capacity and simultaneously NO₂ adsorption turns into the new rate-limiting factor. Generally,
290 it could be deduced that the balance between NO₂ adsorption and PAS formation results in the nonlinear uptrend of
291 γ -values with increasing light intensity.

292 Concentration dependence was considered in this research, and the nitrate formation rates at given conditions
293 were normalized by the corresponding result estimated at 160 mW/cm² (**Figure 4b**). At low concentration (9.20
294 ppm), the formation rate is light-dependent under weak illumination while tends to be steady with increasing light
295 intensity, suggesting excess PAS under strong irradiation. The difference is that, at high concentration (21.45 ppm)
296 the formation rate under strong irradiation was not nearly equal to that under 160 mW/cm², implying sufficient
297 adsorbed NO₂ at relatively high intensity. Generally, higher NO₂ concentrations suggest broader influence scope
298 of illumination. Hence, current serious NO₂ pollution may increase the participation of solar irradiation in the
299 formation of secondary aerosols.

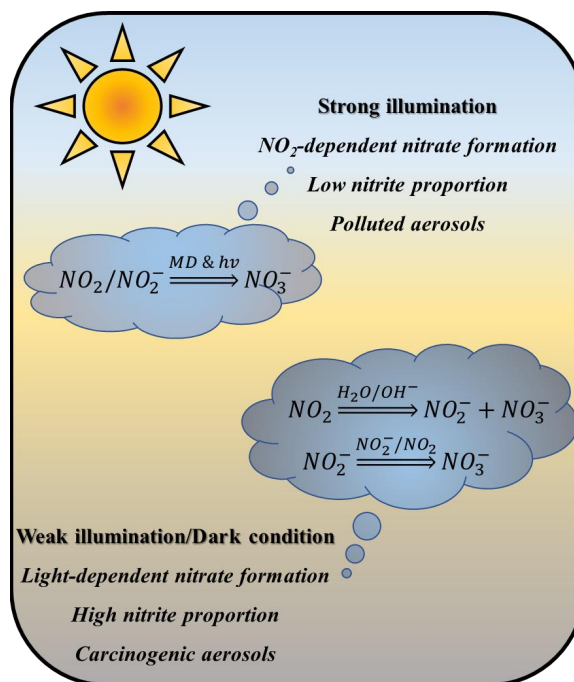


300

301 **Figure 4.** (a) Current densities of TiO₂ under various light intensities (60-120 s). Inset: linear correlation between
302 averaged photocurrent densities and irradiation intensities. (b) Relative formation rate of nitrate as a function of
303 light intensity under different NO₂ concentrations.

304 On the other hand, limited PAS participate in the oxidation of excess NO₂ under weak illumination, and thus
305 make space for the nitrite formation via disproportionation process (R.1). Under strong irradiation, there are still
306 sufficient PAS involved in the nitrite oxidation after the photochemical conversion of limited NO₂ (R.2). For the
307 oxidation of nitrite intermediates, the main promoters are NO₂ and other nitrite species under dark condition or
308 weak illumination (R.3, 4), while turns to be PAS under stronger irradiation (R.5). Hence, nitrite products
309 unevenly decrease with increasing light intensity.

310 As shown by Scheme 1, sunlight influences the formation pathways of nitrate and nitrite aerosols, and finally
311 results in reactions with different features. Mineral dust under weak illumination (or dark condition) and strong
312 irradiation may be covered by different nitrogen compounds: nitrite and nitrate, respectively. Since nitrite may
313 induce cancer risk while nitrate is generally treated as secondary pollutants, we can name them carcinogenic
314 aerosols and polluted aerosols in atmospheric research, respectively.



315

316 **Scheme 1.** Characteristics of the photoinduced heterogeneous reaction of NO₂ on mineral dust (MD) under
317 different illumination conditions.

318 **4 Conclusions and atmospheric implications**

319 Nitrate is dominating atmospheric particulates with the increasing NO_x emissions from expanding urban
320 traffic (Anenberg et al., 2017). Photochemistry has traditionally been considered inapparent during haze events
321 because of the weak sunlight near the ground caused by low visibility (Cheng et al., 2016; Shen et al., 2015;
322 Zhang et al., 2015). However, the nitrate formation on mineral dust is found to be more dependent on weak
323 sunlight, indicating that photochemistry processes are still crucial in heavy haze. Since the NO₂ concentrations in
324 the troposphere are much lower than the simulated levels, authentic dust may be close to achieving its highest
325 uptake capacity in the presence of faint sunlight (Figure 4b). Hence, photoinduced reaction on mineral dust may
326 contribute greatly to secondary aerosols during extreme haze events.

327 Nitrate pollution has got much concern recently, while little attention has been paid to the nitrite burst
328 accompanied by low nitrate concentration. Nitrite may induce adverse health risk for its close association with
329 various cancer cases (Zhang et al., 2018). Compared to the polluted aerosols with high nitrate level, the
330 carcinogenic aerosols with great nitrite concentration may be more harmful to human health. As an intermediate in
331 the photochemistry activities, nitrite appears to be the main product under weak sunlight. The light-dependent



332 negative correlation between nitrate and nitrite highlights illumination an inducing factor in the atmospheric
333 nitrogen cycling.

334 Actually, we discussed the γ -values based on the average experimental results while gave little care to various
335 measurement errors. Compared to the arithmetic mean results, the percentile γ -values estimated by Monte Carlo
336 simulation could be more suitable for modelling studies due to the differences between real atmosphere and the
337 simulated laboratory condition. Furthermore, sensitivity analysis is helpful in explaining the determining factors
338 involved in the assessment of uptake capacity. Generally, statistical simulation brings about more accurate
339 evaluation and provides opportunities to explain the model discrepancy for secondary aerosols.

340

341 **Data availability.** All data are available upon request from the corresponding authors.

342 **Supporting information.** Sections on particle characterization, experimental setup, pretreatment for in-suit
343 DRIFTS test, uptake coefficient estimation, field observations, product observations, detailed reactions in
344 photocatalytic process, photoinduced nitrite oxidation, Raman detection, Sensitivity analysis.

345 **Author contributions.** TW designed the experiments and wrote the paper. YYL and YD contributed to the
346 DRIFTS spectra analysis. YYL and XZF support the field observation. HYC, YQF, MT, and XD assisted the
347 Raman measurements. YY, KJL, SA, AB, and IN performed the PEC tests. LWZ guided the data analysis and
348 paper writing. HBF and JMC provided some experimental facilities. All authors were involved in the discussion.

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