

Sulfate formation during heavy winter haze events and the potential contribution from heterogeneous SO₂+NO₂ reactions in the Yangtze River Delta region, China

Ling Huang¹⁺, Jingyu An²⁺, Bonyoung Koo³, Greg Yarwood⁴, Rusha Yan², Yangjun Wang¹, Cheng Huang^{2*}, Li Li^{1*}

¹School of Environmental and Chemical Engineering, Shanghai University, Shanghai, 200444, China

²State Environmental Protection Key Laboratory of the Cause and Prevention of Urban Air Pollution Complex, Shanghai Academy of Environmental Sciences, Shanghai, 200233, China

³Bay Area Air Quality Management District, San Francisco, 94105, USA

⁴Ramboll, Novato, California, 95995, USA

⁺These two authors contributed equally to this work.

Correspondence to: Li Li (lily@shu.edu.cn) and Cheng Huang (huangc@saes.sh.cn)

Abstract. Rapid sulfate formation is recognized as key characteristics of severe winter haze in China. However, air quality models tend to underestimate sulfate formation during heavy haze periods and heterogeneous formation pathways have been proposed as promising mechanisms to reduce gaps between observation and model simulation. In this study, we implemented a reactive SO₂ uptake mechanism through the SO₂ + NO₂ heterogeneous reactions in the Comprehensive Air Quality Model with extensions (CAMx) to improve simulation of sulfate formation in the Yangtze River Delta (YRD) region. Parameterization of the SO₂ + NO₂ heterogeneous reactions is based on observations in Beijing and considered both impact of relative humidity and aerosol pH on sulfate formation. Ammonia is reported to be critical for the formation of secondary inorganic aerosols. Estimation of ammonia emissions is usually associated with large uncertainties and model tends to underestimate ammonia concentrations substantially. Sensitivity tests were conducted to evaluate the influence of the SO₂ + NO₂ heterogeneous reactions as well as ammonia emissions on modelled sulfate concentrations during a period with several heavy haze episodes in the YRD region. Base case model results show large underestimation of sulfate concentrations by 36 % under polluted conditions in the YRD region. Adding the SO₂ + NO₂ heterogeneous reactions or doubling ammonia emissions alone leads to slight model improvement (~6 %) on simulated sulfate concentrations in the YRD region. However, model performance significantly improved when both the SO₂ + NO₂ heterogeneous reactions and doubled ammonia emissions were included in the simulation: predicted sulfate concentrations during polluted periods increased from 23.1 µg m⁻³ in the base scenario to 29.1 µg m⁻³ (representing an increase of 26 %). Aerosol pH is crucial for the SO₂ + NO₂ heterogeneous reactions and our calculated aerosol pH is always acidic and increased by 0.7 with doubled ammonia emissions. Modelling results also show that this reactive SO₂ uptake mechanism enhanced sulfate simulations by 1 to 5 µg m⁻³ for the majority of eastern and central part of China, with more than 20 µg m⁻³ increase of sulfate concentrations over the

north-eastern plain. These findings suggest that the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions could be potentially important for sulfate formation in the YRD region as well as other parts of China. Further studies are needed to constrain the uncertainties associated with the parameterization of the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions based on local data as well as to evaluate this mechanism in other regions. In addition, ammonia emissions were found to be a key driving variable of the spatial patterns of sulfate enhancement due to the new pathway. Substantial efforts are needed to improve the accuracy of ammonia emissions inventory.

1 Introduction

Rapid sulfate (SO_4^{2-}) formation has been reported to be key characteristics of severe winter haze in China. However, most air quality models tend to underestimate sulfate formation during severe winter haze episodes in China because standard SO_2 oxidation pathways, including gas-phase chemistry (i.e. oxidized by hydroxyl radical OH) and aqueous-phase chemistry (i.e. oxidized by ozone (O_3), hydrogen peroxide (H_2O_2)) are suppressed by weak photochemical activity and low ozone concentrations (Quan et al., 2014)). Meanwhile, analysis of severe haze events in China show enhanced secondary inorganic aerosols, especially sulfate concentrations. For example, Quan et al. (2014) found that observed sulfate accounted for 13 % of $\text{PM}_{2.5}$ (particulate matter with dynamic equivalent diameter less than $2.5 \mu\text{m}$) on normal clean days and increased to 25 % on haze days during the infamous 2013 January Beijing haze period. For the same haze episode, Cheng et al. (2016) used concentration ratios of sulfate to sulfur dioxide ($[\text{SO}_4^{2-}]/[\text{SO}_2]$) to diagnose sulfate production rate; this ratio increased with $\text{PM}_{2.5}$ levels and was 6 times higher under the most polluted conditions than normal conditions. Most current air quality models (e.g. CMAQ, GEOS-Chem, WRF-Chem, CAMx), which only include the traditional gaseous- or aqueous-phase mechanisms for sulfate formation, do not show very good model performances of sulfate concentrations against observations during haze periods in China (Wang et al., 2014; Zheng et al., 2015; Gao et al., 2016a, 2016b; Li et al., 2015). The under-prediction of sulfate concentrations could be related to uncertainties of the emissions inventory, bias of simulated meteorological fields, and/or some missing sulfate formation mechanisms that are not included in the current models.

Heterogeneous sulfate production chemistry has been proposed by several studies to explain the high concentrations and rapid formation of sulfate during haze episode in China (e.g. He et al., 2014; Wang et al., 2014; Zheng et al., 2015; Wang et al., 2016; Cheng et al., 2016; Guo et al., 2017). He et al. (2014) suggested a synergistic effect between NO_2 and SO_2 on the surface of mineral dust (i.e. mineral oxides) as an important source of sulfate in China and emphasized the essential role of O_2 involved in this process. More generally, heterogeneous loss of SO_2 on aerosol surfaces (not limited to mineral dust) or deliquescent aerosols is discussed by many studies, although the exact underlying mechanism is still unknown (e.g. Wang et al., 2014; Zheng et al., 2015). For this kind of heterogeneous reactions, the sulfate production rate has been parameterized as a pseudo first-order reaction with respect to the gaseous SO_2 concentration with the SO_2 reactive uptake coefficient (γ) on aerosol surfaces being the key parameter. This uptake coefficient, representing the probability that a SO_2 gas molecule colliding with an aerosol surface results in sulfate formation, is reported to be heavily dependent on relative humidity (RH)

(Zheng et al., 2015; Wang et al., 2016). Parameterized reactive uptake of SO_2 has been implemented in several current air quality models, including GEOS-Chem, WRF-Chem, CMAQ and CAMx, and generally improved model performance of sulfate concentrations during haze episodes in China (e.g. Wang et al., 2014; Zheng et al., 2015; Gao et al., 2016b). Two more recent papers, Wang et al. (2016) and Cheng et al. (2016) further suggested that reaction between NO_2 and SO_2 in aerosol water may contribute substantially to sulfate formation during haze events in China. Both studies emphasized the importance of higher aerosol pH (5.4–6.2 reported by Cheng et al. (2016) and 6.0–7.6 by Wang et al. (2016)) sustained by abundant gas-phase ammonia (NH_3) during haze periods being an essential precondition for this mechanism. However, the near-neutralized aerosol pH that facilitates SO_2 oxidation by NO_2 is questioned by Guo et al. (2017) who concluded from a thermodynamic analysis that aerosol pH was always acidic (4.5–5) regardless of the ambient NH_3 concentrations and that the NO_2 -mediated oxidation of SO_2 was unlikely to be important in China or any other region of the world. Guo et al. (2017) pointed out that within low pH ranges (up to 4.5), SO_2 oxidation catalyzed by transition metal (i.e. Fe(III) and Mn(II)) might become a dominant sulfate formation pathway in aerosol water and suggested it as an alternative to $\text{SO}_2 + \text{NO}_2$ reactive uptake as being a potential sulfate contributor under haze conditions. Similar conclusion is also made from a most recent work by Shao et al. (2019), who implemented four heterogeneous sulfate formation mechanisms in GEOS-Chem and assessed model performance using sulfate oxygen isotopes data in Beijing, who found that SO_2 oxidation catalyzed by transition metal ion (TMI) to be the dominant sulfate formation mechanism. On the contrary, another slightly earlier study by Ye et al. (2018) concluded SO_2 oxidation by H_2O_2 was the dominant pathway based on observations of atmospheric H_2O_2 concentrations in Beijing. Song et al. (2018b) suggested the heterogeneous hydroxymethanesulfonate (HMS) chemistry being a potentially important contributor to heavy haze pollution in northern China. Hung et al. (2018) reported the interfacial SO_2 oxidation on the surface of aqueous micro-droplets as a potential pathway to explain fast conversion of SO_2 to sulfate.

To investigate whether the $\text{SO}_2 + \text{NO}_2$ reactions in aerosol water could help better predict the enhanced sulfate formation during haze periods in the Yangtze River Delta (YRD) region, we implemented a parameterized $\text{SO}_2 + \text{NO}_2$ reactive uptake mechanism in the Comprehensive Air Quality Model with Extensions (CAMx), which is a widely used air quality model in China (e.g. Wang et al., 2009; Huang et al., 2012; Li et al., 2013, 2015; Jia et al., 2017; etc.). Our parameterization specifically incorporated RH and aerosol pH dependencies derived from measurement data during the 2015 Beijing haze event (Wang et al., 2016). Although the RH dependency of the SO_2 uptake rate has already been implemented in previous studies (e.g. Zheng et al., 2015; Wang et al., 2014), the effect of aerosol pH has not been explicitly included in most of the previous modelling studies, except for a most recent study by Shao et al (2019), who also considered aerosol pH in their model parameterization.. While most of the previous studies were trying to improve model predictions in the northern part of China, especially the Beijing-Tianjin-Hebei region (e.g. Gao et al., 2016b; Zheng et al., 2015), this work is one of the few studies that focus on the Yangtze River Delta region, which has also suffered from severe haze problems in recent years due to urban expansion and industrialization (e.g. Li et al., 2011; M. Wang et al., 2015; Xu et al., 2016; Ming et al., 2017). In addition to the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions, we also investigated model sensitivities to ammonia emissions, which

have been reported to be crucial for the formation of secondary inorganic aerosols and large uncertainties exist with current ammonia emission inventory (Huang et al., 2011; Fu et al., 2013).

2 Methodology

2.1 Current sulfate formation pathways in CAMx

- 5 In this study, CAMx version 6.40 (Ramboll Environ, 2016) was used as the base model to simulate sulfate formation. Table 1 lists the sulfate formation pathways that are currently considered implemented in standard CAMx source code. In addition to the traditional SO₂ oxidation by OH in the gas phase and O₃, H₂O₂, and O₂ (catalyzed by Fe(III)/Mn(II)) in cloud droplets, sulfate formation through reactions with methyl hydroperoxide and other organic hydroperoxides (MHP) as well as peracetic and other organic peracids (PAA) in the aqueous phase is also included. For heterogeneous formation pathway, the SO₂ + NO₂ reaction is currently considered and implemented as pseudo gas phase reaction with the rate parameterization based on results from Zheng et al (2015), where the key parameter (i.e. gamma) is bounded between a lower and upper limit and changes linearly in response to RH. This relatively simple parameterization of SO₂ + NO₂ heterogeneous reaction has been included in many previous studies, e.g. Y. Wang et al. (2014), B. Zheng et al. (2015), etc.

2.2 SO₂ + NO₂ mechanism in CAMx

- 15 In this study, we implemented the SO₂ + NO₂ reactive uptake mechanism in CAMx version 6.40 (Ramboll Environ, 2016) as a pseudo gas-phase reaction:



Since the vapor pressure of sulfuric acid is very low, we assumed all sulfuric acid partitions to the aerosol phase. The rate constant k_{het} is related to the reactive uptake coefficient γ for SO₂ as follows:

$$\frac{d[SO_4^{2-}]}{dt} = k_{het}[NO_2(g)][SO_2(g)] = \frac{1}{4}\gamma\bar{C}S[SO_2(g)] \quad (2)$$

- 20 where \bar{C} is the mean molecular speed (m/s), and S is the aerosol surface area concentration (m²/m³). Based on the observations during the Chinese haze events (Wang et al., 2016), this uptake coefficient γ depends on aerosol pH, RH, and NO₂ concentration. Therefore, we assumed a functional form of γ as the product of each of these dependencies:

$$\gamma = 4k_0d_f[NO_2(g)] \quad (3)$$

where k_0 (ppm⁻¹) is the RH-dependent parameter; NO₂(g) is the NO₂ gas concentration; d_f is the pH-dependent distribution factor of SO₂, i.e. the ratio of SO₂ concentration in the aqueous-phase to the gaseous-phase and is calculated as Eq. (4) in the model:

$$d_f = \frac{[S(IV)(aq)]}{[SO_2(g)]} = H_{eff}RTw_L \quad (4)$$

where H_{eff} is the effective Henry's law constant of SO_2 (M atm^{-1}), R is the universal gas constant ($\text{L atm mol}^{-1} \text{K}^{-1}$), T is air temperature (K) and w_L is the aerosol water content ($\mu\text{g m}^{-3}$). We used the data in Table S2 and S5 of Wang et al. (2016) to back calculate the RH dependency of k_0 under clean (observed sulfate concentration less than $10 \mu\text{g m}^{-3}$), transition (sulfate between 10 and $20 \mu\text{g m}^{-3}$), and polluted (sulfate more than $20 \mu\text{g m}^{-3}$) conditions during Beijing 2015 episodes. The reactive uptake coefficient γ derived by Wang et al. (2016) contains an assumption that all sulfate formation was due to the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions due to low photochemical reactivity and low Fe/Mn concentrations during the measurement period. Thus we started with an upper bound estimate of the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions. Aerosol pH was calculated using the ISORROPIA thermodynamic equilibrium model implemented in CAMx assuming a metastable aerosol liquid phase which is an appropriate assumption for most ambient conditions including the Chinese haze events (Guo et al. 2017). Wang et al (2016) only reported NO_x (not NO_2) concentrations in Beijing during the 2015 haze event. We simply assumed a NO_2/NO_x ratio of 0.5. Inserting NO_2 concentrations, γ values from Wang et al. (2016), and calculated aerosol pH from ISORROPIA into Eq. 3, we obtained the expression of k_0 depending upon RH as follows (parameters for k_0 calculation is shown in Table S1):

$$k_0 = \begin{cases} \text{RH} < 21\%: & 199.25 \\ 21\% \leq \text{RH} < 41\%: & (284.22 - 199.25) \times (\text{RH} - 21\%) / (41\% - 21\%) + 199.25 \\ 41\% \leq \text{RH} < 56\%: & (322.16 - 284.22) \times (\text{RH} - 41\%) / (56\% - 41\%) + 284.22 \\ \text{RH} \geq 56\%: & 332.16 \end{cases} \quad (5)$$

Due to lack of observation data at high RH values, as a conservative assumption we set a constant k_0 value when RH increases from 56% and up. This would lead to underestimated sulfate formation due to the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions at high RH values, which is a favorable condition for the heterogeneous sulfate production. In addition, the differences of aerosol hygroscopicity in Beijing vs. Shanghai could add more uncertainties in the dependency of k_0 on RH. Reported values of hygroscopicity parameter κ were 0.25~0.31 for Shanghai (Ye et al., 2011; 2013), which are higher than values reported for Beijing (0.14~0.24; Massling et al., 2009). Our conservative assumption that the heterogeneous reaction rate does not increase with RH above 56% most likely tends to offset the upper bound estimate of the γ values derived by Wang et al. (2016). The rate constant k_{het} of $\text{SO}_2 + \text{NO}_2$ is formulated as:

$$k_{\text{het}} = k_0 d_f \bar{C} S \quad (6)$$

SO_2 lifetime (in hr) associated with the $\text{SO}_2 + \text{NO}_2$ reactive uptake mechanism is calculated as:

$$\text{SO}_2 \text{ lifetime} = \frac{1}{k_{\text{het}} [\text{NO}_2(g)]} \quad (7)$$

Figure 1 shows the SO_2 lifetime as a function of aerosol pH for clean, transition, and polluted conditions, with other variables kept constant. The SO_2 lifetime shortens as aerosol pH becomes more neutralized, indicating faster conversion of SO_2 to sulfate by $\text{SO}_2 + \text{NO}_2$ reactive uptake on aerosol. For pH within 2 to 7, one unit increase in aerosol pH shortens SO_2 lifetime by about one order of magnitude. The blue, orange, and red symbols in Figure 1 correspond to the clean, transition,

and polluted conditions during Beijing 2015 based on data in Table S1. As shown in Figure 1, the aerosol pH values calculated by ISORROPIA are 5.5 (for clean conditions) and 4.1-4.2 (for transition and polluted conditions), all lower than the values (7.6) reported by Wang et al. (2016). As noted by Guo et al. (2017), it is important to make a consistent assumption for aerosol state (i.e., metastable) in deriving and implementing the parameterization for reactive uptake. A most recent paper by Song et al (2018a) identified coding errors with the ISORROPIA model, which resulted unrealistic pH values of 7.7 using the standard ISORROPIA model with the stable state assumption in previous studies. Nevertheless, our results are not compromised by this coding error because the metastable assumption was chosen in our ISORROPIA calculation.

2.3 Model configuration

Two versions of the Comprehensive Air Quality Model with extensions (CAMx) modified based on the original version 6.40 (Ramboll Environ, 2016) were used in this study: one with the SO₂ + NO₂ heterogeneous reactions (described in Section 2.1) and one without (forcing k_{het} equals to zero). The modeling domain consists of three nested grids (Figure 2): the outer 36 km domain (D01) covers most of China, Japan, Korea, parts of India, and southeast Asia; the 12 km domain (D02) covers eastern China and the inner 4 km domain (D03) covers Shanghai, Jiangsu province, Zhejiang province, Anhui province, and parts of surrounding provinces, together referred as the Yangtze River Delta (YRD) region. Meteorological fields were based on simulation results from the Weather Research and Forecasting (WRF) model (version 3.7) driven by the National Centers for Environmental Prediction (NCEP)/National Center for Atmospheric Research (NCAR) Operational Global Analysis data (<http://dss.ucar.edu/datasets/ds083.2/>). Details of the WRF configurations can be found in previous studies (Liu et al., 2018). Boundary conditions for D01 were generated from the Model for OZone And Related chemical Tracers (MOZART) global chemistry model (Emmons et al., 2010). The Carbon Bond 6 (CB6) mechanism (Yarwood et al., 2010) was used for the gas phase chemistry and the static two-mode coarse/fine (CF) scheme was used to represent particle size distribution. The Zhang dry deposition (Zhang et al. 2003) and default wet deposition scheme was used to for removal processes. Anthropogenic emissions for areas outside the YRD region were from the Multi-resolution Emission Inventory for China (MEIC, <http://www.meicmodel.org/>). For emissions within the YRD region, an YRD-specific emission inventory (Huang et al., 2011; Li et al., 2011) was updated to year 2014 and utilized in this study. This YRD-specific emission inventory includes emissions from sources of combustion, industry, mobile and residential. Primary sulfate emissions over the 4km domain are estimated to be 994 tons day⁻¹ for December 2013 (accounting for 14.8% of primary PM_{2.5}) with dense emissions from Shanghai and southern Jiangsu province (see Figure S1 for spatial distribution). At the SAES site, primary sulfate emissions were estimated to 757 kg per month (only accounting for 1.0% of primary PM_{2.5}). Biogenic emissions were simulated using the Model of Emissions of Gases and Aerosols from Nature (MEGAN, version 2.1, Guenther et al. 2012) based on the WRF simulation results. The modeling episode is December 2013, during which several heavy haze events with hourly PM_{2.5} concentration higher than 500 µg m⁻³ were observed in the YRD region.

Four simulations with identical model configuration and input data including meteorology, initial/boundary conditions, and emission inventory (except ammonia emissions) were conducted using the above two different CAMx versions:

- noHet (base case): simulation based on CAMx version without the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions (this is also our base case). Note that this CAMx version differs from the distributed CAMx v6.40 in that we removed the original heterogeneous sulfate formation reaction with NO_2 which only included a simple parameterization based on RH (ref. reaction No.7 in Table 1) in the distributed version. This is done on purpose to quantify the influence of the newly parameterized $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions on sulfate formation.
- Het: simulation based on CAMx with the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions. Other model configurations were identical to scenario noHet.
- noHet_2NH₃: CAMx version and model configurations were same as scenario noHet except ammonia emissions were doubled for the 4 km domain.
- Het_2NH₃: CAMx version and model configurations were same as scenario Het but ammonia emissions were doubled for the 4 km domain.

We first ran CAMx for 36 km/12 km domains with two-way nested; for the 4 km domain, we used boundary conditions extracted from the 12 km model outputs and conducted the above four scenarios. Fourteen vertical layers were used extending from the surface to 100 mb. In addition to default CAMx outputs, we modified the source code to generate additional diagnostic variables (e.g. aerosol pH, RH, and k_{het}) to evaluate the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions.

2.4 Observations

Hourly observations of ozone, SO_2 , NO_2 , $\text{PM}_{2.5}$ and its components including sulfate, nitrate, ammonium, organic carbon (OC), and elemental carbon (EC) are available between 1 December and 29 December 2013 at a monitor site located at the center of the urban area of Shanghai (referred as SAES site, 31.1695°N, 121.4305°E, Figure 3). Hourly $\text{PM}_{2.5}$ observations are also available at another 23 monitor sites across the YRD region (Figure 3; see locations in Table S2). During this period, YRD region experienced relative clean days as well as several heavy haze episodes with peak $\text{PM}_{2.5}$ exceeding $600 \mu\text{g m}^{-3}$ during a most heavily polluted period of December 5th to 7th. At the SAES site, maximum hourly $\text{PM}_{2.5}$ concentration reached $540.3 \mu\text{g m}^{-3}$ on December 6th with a monthly average of $118.7 \mu\text{g m}^{-3}$. We followed the method in Wang et al (2016) to divide the period into clean (observed sulfate $<10 \mu\text{g m}^{-3}$), transition ($10\text{--}20 \mu\text{g m}^{-3}$), and polluted ($>20 \mu\text{g m}^{-3}$) periods based on observed hourly sulfate concentration at the SAES site. Compared with clean period, all PM species increased by more than 3 times (sulfate, nitrate and ammonium (SNA) increase by more than 5 times) during polluted period as indicated by the enhancement ratio (calculated as the ratio of average concentrations during the polluted period divided by those during the clean period). In terms of fraction of $\text{PM}_{2.5}$, SNA increased from 44 % during clean period to 69 % during polluted period while carbonaceous aerosols (OC and EC) decreased from 32 % to 24 %. This is consistent with the commonly observed characteristics of winter haze periods in China reported by many previous studies (e.g. Wang et al., 2014; Zheng et al., 2015b; Cheng et al., 2016) that SNA is playing a more important role during the heavy haze periods.

Average sulfate concentration of clean, transition and polluted periods was 6.7, 14.2, and 36.1 $\mu\text{g m}^{-3}$, respectively, accounting for 17-23 % of $\text{PM}_{2.5}$ (Figure S2).

Observations of ambient ammonia concentrations are also available at the SAES site; however, the quality of measurements is questionable. Therefore, we used ammonia observations from a similar urban site nearby (referred as FDU site, ~15 km north from the SAES site, 31.3005°N, 120.9778°E, Figure 3) for analysis in this study. Observations at the FDU site have been discussed by S. Wang et al. (2015) and demonstrated data reliability. Diurnal NH_3 concentrations at the FDU site during our modeling period showed a weak bimodal pattern with an average of 7.3 ppb (ranging 1.6–25 ppb) during this period (Figure S3). This two-peak diurnal variation is caused by vehicle emissions and evolution of the boundary layer (S. Wang et al. 2015). In summary, observations for gases species (except NH_3) and PM species at the SAES site and NH_3 at the FDU site were used for model validation in this study.

2.5 Statistical metrics for model validation

For WRF and CAMx model performance valuation, mean bias (MB), normalized mean bias (NMB), and index of agreement (IOA) were used in this study. Calculations of these selected metrics are shown below:

$$MB = \frac{1}{N} \sum (P_j - O_j) \quad (8)$$

$$NMB = \frac{\sum (P_j - O_j)}{\sum O_j} \times 100 \quad (9)$$

$$IOA = 1 - \frac{\sum (P_j - O_j)^2}{\sum (|P_j - \bar{O}| + |O_j - \bar{O}|)^2} \quad (10)$$

where P_j and O_j are predicted and observed hourly concentrations or values, respectively. N is the number of paired model and observation data. \bar{O} is the average concentration/value of observations. IOA ranges from 0 to 1 with 1 indicating perfect agreement between model and observation.

3 Results and discussions

3.1 Model evaluation

3.1.1 WRF results evaluation

Model performance of WRF results is generally acceptable in this study. Table S3 summarizes the meteorological performance statistics of WRF during December 2013 at Pudong and Hongqiao airport stations in Shanghai (Figure 3). Temperature and relative humidity were well reproduced with NMB and NME within 37% and 41%, respectively; IOA values are above 0.8. Wind speed is overestimated with a MB of 1.5 m s^{-1} at Pudong and 0.5 m s^{-1} at Hongqiao station; NMB of predicted wind direction at the two stations is -36% and -27%, respectively. Comparisons of hourly observed and simulated relative humidity, wind speed and temperature at these two stations suggest reasonable model results in terms of

temporal variations (Figure S4). Overall, the WRF simulated results are acceptable to be used in subsequent CAMx simulations.

3.1.2 CAMx base scenario (noHet) evaluation

Figure 4 depicts the time series of simulated and observed concentrations for sulfate and PM_{2.5} during 1 to 29 December 2013 at SAES site (see Figure S5 in Supplemental Information for other species). Overall, the model is successful in capturing the temporal variations of ozone and PM species with IOA values above 0.5 (Table S4). Nevertheless, model tends to systematically underestimate all gaseous and PM species with NMB values ranging from -5% for NO₂ to -68% for NH₃. This could be partially explained by the higher simulated wind speeds compared with observed values, especially at Pudong station where the observed average wind speed during the modeling period was 4.5 m s⁻¹ while simulated wind speed was 6.0 m s⁻¹, representing an overprediction by 33%. For sulfate, the model captured the day-to-day sulfate variations reasonably well with an overall MB of -2.8 µg m⁻³ and IOA of 0.80. For clean and transition periods, model showed slight over-prediction with MB of 1.1 and 0.5 µg m⁻³ (Table S5). However, during polluted period when observed sulfate concentrations are higher than 20 µg m⁻³, model significantly underestimated sulfate formation with a MB of -13.0 µg m⁻³ (NMB of -36 %). Observed maximum sulfate concentration reached 93.4 µg m⁻³ but model only predicted 52.2 µg m⁻³. Nitrate and ammonium concentrations were also underestimated by 20 % on average and exacerbated to more than 40 % during polluted periods. For carbonaceous aerosols, elemental carbon (EC) was underestimated by 32 % while organic carbon (OC) exhibited even more underestimation of almost 50 %. Underestimation of OC is usually associated with underestimation of secondary organic aerosols (SOA). Discussion of OC under-prediction is beyond the scope of current work and will be addressed in future studies. Results of the four CAMx simulations in this study showed negligible changes in predicted EC/OC concentrations and thus are excluded in the following discussions.

Figure 5 depicts the averaged PM_{2.5} during the modeling episode over the YRD region with observations at 24 monitoring sites. Observed PM_{2.5} concentrations generally showed a decreasing trend from north to south of the YRD region, which was well captured by the model. For sites located in southern Jiangsu and southern Zhejiang province, the model showed favorable agreement with the observations. Underestimations existed for sites located in the northern part of Jiangsu and Zhejiang province. MB across all 24 monitoring sites ranged from as low as -90.4 µg m⁻³ (site in north Jiangsu province) to slight overestimation of 11.4 µg m⁻³ (site in south Zhejiang province); corresponding NMB ranged from -46 % to 16 % (Table S2).

3.2 Simulated sulfate concentrations at SAES site

Four scenarios – noHet, Het, noHet_2NH₃ and Het_2NH₃ were conducted to evaluate the impact of the SO₂ + NO₂ heterogeneous reactions and ammonia emissions on sulfate simulation. We first analyzed the modeled sulfate results at the SAES site; then we discussed the spatial patterns over the YRD region. Similar discussions of nitrate, ammonium and PM_{2.5} are included in the supplemental information. Table 2 shows the average sulfate concentration for different scenarios by

clean, transition, and polluted periods; corresponding scatter plots are shown in Figure 6. A complete summary of statistical metrics for each scenario/period is presented in Table S5.

Impact of SO₂ + NO₂ heterogeneous reactions (noHet vs. Het)

As shown in Figure 6, simulated sulfate concentrations compared well with observations under clean and transition conditions in the noHet scenario with over-prediction by 16 % and 4 %, respectively. By contrast, large under-prediction of sulfate concentration existed during polluted periods (MB of -13.0 $\mu\text{g m}^{-3}$, NMB of -36 %). Adding the SO₂ + NO₂ heterogeneous reactions showed small enhancement on sulfate formation, reducing the overall NMB from -16 % to -12 %. If only polluted periods are considered, simulated sulfate concentrations increased from 23.1 to 24.6 $\mu\text{g m}^{-3}$ with the heterogeneous reactions, corresponding to an increase by 6.5 %. Thus even with the SO₂ + NO₂ heterogeneous reactions, model was still under-predicting sulfate concentrations on heavy haze days with a NMB of -32 %. This is because aerosol pH was always acidic (pH < 3; this will be discussed in the following section) and the SO₂ + NO₂ heterogeneous reactions were not being appreciable within this pH range (Figure 1). Model performances for clean and transition periods were slightly compromised with the SO₂ + NO₂ heterogeneous reactions since the base scenario was already overestimating sulfate concentrations.

Impact of NH₃ emissions (noHet vs. noHet_2NH₃)

Being the dominant base gas in the atmosphere, ammonia plays an essential role in the formation of secondary inorganic aerosols and estimation of ammonia emissions is usually associated with large uncertainties (e.g. Huang et al., 2011; Fu et al., 2013). With the base case ammonia emissions, NH₃ concentration was under-predicted by 3.0 ppb (NMB of -60 %). With doubled ammonia emissions, ammonia concentration was over-predicted by 1.7 ppb with NMB of 34 % but the MB of the total ammonia (NH₃ + ammonium) concentrations were reduced from -6.9 $\mu\text{g m}^{-3}$ (NMB of -36%) in the base case scenario to -1.9 $\mu\text{g m}^{-3}$ (NMB of -10%). NMB of sulfate concentrations during polluted period is -32 %, which is similar to the enhancement caused by that of the SO₂ + NO₂ heterogeneous reactions. Clearly, doubling ammonia emissions is not enough to close the gap between observed and simulated sulfate concentrations during heavy haze periods. We performed additional sensitivity tests with even higher ammonia emissions and found that 10 times ammonia emissions would be needed to achieve an average sulfate concentration (33.2 $\mu\text{g m}^{-3}$) that is comparable with observation (36.1 $\mu\text{g m}^{-3}$) under polluted conditions (with no heterogeneous reactions). However, in that case, model performance of ammonia is significantly compromised with over-prediction by 32.3 ppb. These results indicate that the uncertainties associated with the ammonia emissions are not enough to fully explain the under-prediction of sulfate formation during heavy haze periods in the YRD region.

Impact of both (noHet vs. Het_2NH₃)

A fourth scenario (Het_2NH₃) with the SO₂ + NO₂ heterogeneous reactions as well as doubled ammonia emissions gave the best model performance of sulfate concentrations with an overall MB of -0.2 $\mu\text{g m}^{-3}$ (NMB of -1 %, Figure 6). During polluted periods, average sulfate concentration was predicted to be 29.1 $\mu\text{g m}^{-3}$ (representing an increase of 26% from the base case) and NMB was reduced from -36 % in the base scenario to -19 % in the Het_2NH₃ scenario. Maximum sulfate

concentration simulated under scenario Het_2NH₃ was 97.2 $\mu\text{g m}^{-3}$, which compared well with the observed maximum of 93.4 $\mu\text{g m}^{-3}$ at the SAES site. With doubled ammonia emissions, the heterogeneous reactions were playing an increasing important role in sulfate formation by boosting average sulfate concentrations from 24.5 (noHet_2NH₃) to 29.1 $\mu\text{g m}^{-3}$ (Het_2NH₃) under polluted conditions, representing an increase by 19 %. This is because with more ammonia available, aerosol pH was elevated by ~ 0.7 pushing it closer towards the actual pH (as discussed more in section 3.3) and the rate of the heterogeneous reactions is positively correlated with aerosol pH (Figure 1), therefore leading to the best model performances from the Het_2NH₃ scenario. These results indicate that the SO₂ + NO₂ heterogeneous reactions as well as sufficient ammonia emissions are both needed to greatly improve model simulation of sulfate formation under polluted conditions. However, it is to mention that model performance under clean and transition periods got compromised most under scenario Het_2NH₃.

Figure S6 shows a Q-Q plot of modeled versus observed sulfate concentrations for the four scenarios. Underestimations of sulfate concentrations become noticeable around 35 $\mu\text{g m}^{-3}$ in all scenarios and between 35 to 55 $\mu\text{g m}^{-3}$, there appears to be a systematic low bias in predicted sulfate concentrations that neither doubled ammonia emissions nor the heterogeneous reactions or both could stimulate notable sulfate formation. Scenario Het_2NH₃ gives the best model performance with an overall MB of -0.2 $\mu\text{g m}^{-3}$ but still underpredicts sulfate formation under heavy haze periods by -19 %. This could be related to still biased ammonia emissions, less direct emissions of sulfate and/or SO₂, and/or missing of other sulfate formation pathways that needs further investigation. For example, Shao et al. (2019) included heterogeneous sulfate formation via oxidations by O₃, H₂O₂, and Fe(III)/Mn(II), in addition to the aqueous phase reactions and concluded that the metal-catalyzed reactions dominated the heterogeneous sulfate formation. These heterogeneous reactions were not included in the current study and could lead to some underestimated of sulfate formation. As mentioned above, the parameterization of the k_0 values is relatively conservative at high RH conditions, which are favorable for sulfate formation. In addition, reported aerosol hygroscopicity Bias in meteorology could also play some roles here as we are seeing systematically underestimation of all gaseous and PM species. Another explanation is that the SO₂ + NO₂ heterogeneous reactions implemented in this study were parameterized based on observations in Beijing but the simulation is performed over the YRD region. It would be ideal to use local observations for model parameterization in future studies.

Sulfate formation budget

To gain a closer look at the sulfate formation via different pathways (e.g. gas phase vs. aqueous phase vs. heterogeneous phase, Table 1), we constructed a sulfate formation budget in a similar manner as Shao et al. (2019). Figure 7 shows the relative contribution of primary sulfate emissions as well as individual sulfate formation pathway to the total sulfate concentrations at the SAES site under different conditions. Overall, primary sulfate emissions and secondary formation accounted for half of the total sulfate concentrations. Of the secondary sulfate, gas-phase reactions always dominated secondary sulfate formation, with relatively consistent contribution around 38~39% under different conditions. As pollution developed, contribution from secondary formation exceeded that of primary emissions, accounting for 60% of total sulfate abundances under polluted conditions. In contrast to the relatively consistent contribution from the gas-phase formation, both

aqueous and heterogeneous sulfate formation doubled from clean to polluted periods, with relative contribution increased from 4.1% to 9.4% for the former and from 5.0% to 12.6% for the latter.

If we exclude the contribution of primary sulfate emissions (i.e. smaller pie chart in Figure 7), the absolute sulfate formation via the gas-phase reactions more than doubled from clean ($1.59 \mu\text{g m}^{-3}$) to polluted ($3.61 \mu\text{g m}^{-3}$) periods; however, the relative contribution from gas-phase formation among all formation pathways dropped from 80.9% to 63.3% as pollution developed. Sulfate formation from all aqueous phase reactions increased from $0.17 \mu\text{g m}^{-3}$ under clean conditions to $0.89 \mu\text{g m}^{-3}$ under polluted conditions, corresponding to an increase of relative contribution from 8.6% to 15.6%. Under all conditions, aqueous oxidation due to MHP and PAA is negligible, with less than 1% of sulfate contribution. The rest three aqueous pathways in turn dominated aqueous sulfate formation depending on the specific condition. For instance, under clean conditions, oxidation by O_3 was the dominant aqueous contributor (accounting for 5.4% of all sulfate formation pathways) but ignorable ($<1\%$) under polluted conditions. While modeled SO_2 concentrations increased from $33.2 \mu\text{g m}^{-3}$ to $53.5 \mu\text{g m}^{-3}$ as pollution developed, simulated O_3 concentrations dropped by almost half from 8.7 ppb ($\sim 18.7 \mu\text{g m}^{-3}$) under clean conditions to 5.2 ppb ($\sim 11 \mu\text{g m}^{-3}$) under polluted conditions, leading to reduced sulfate formation from aqueous oxidation by O_3 under more severe haze. Predicted O_3 concentrations in this study are much higher than the values (~ 1 ppb) assumed by Cheng et al. (2016) and Wang et al. (2016) but are comparable to values reported by Shao et al (2019; 9 ppb) for a haze episode in Beijing.

Sulfate formation associated with H_2O_2 and Fe(III)/Mn(II) showed the opposite trend: the H_2O_2 pathway only contributed 1.4% ($0.03 \mu\text{g m}^{-3}$) of total sulfate formation under clean conditions and increased to 5.6% ($0.12 \mu\text{g m}^{-3}$) under polluted conditions, representing an increase by a factor of 3. Predicted H_2O_2 concentrations at the SAES site was 0.03 ppb on average and maximum value could reach 0.29 ppb. These values are slightly lower than the values observed in Beijing (average around 0.05 ppb and maximum of 0.90 ppb) by Ye et al. (2018) but are comparable in term of the magnitude. However, without any H_2O_2 observations in Shanghai, it would be inappropriate to conclude whether model is over- or under-predicting H_2O_2 levels in Shanghai. Based on our current results, H_2O_2 oxidation is not the dominant contributor to sulfate formation during our study period.

Oxidation pathway involving Fe(III)/Mn(II) also contributed more to sulfate formation as polluted developed. Under polluted conditions, Fe(III)/Mn(II) catalyzed sulfate oxidation is the dominant aqueous formation pathway, accounting for 8.4% ($0.48 \mu\text{g m}^{-3}$) of secondary sulfate formation. CAMx estimates the Fe(III) and Mn(II) concentrations by assuming a constant mass fraction (3.355% for Fe(III) and 1.15% for Mn(II) based on the dust and primary $\text{PM}_{2.5}$ concentrations. A value of 10% for Fe (III) and 50% for Mn (II) was set to represent the soluble fraction in the cloud water. 10% of Fe(III) during the day and 90% of Fe(III) during the night and all Mn(II) were assumed to be in the oxidized ionic state. Based on these assumptions, modeled soluble concentrations during December 2013 was $1.51 \pm 1.68 \mu\text{M}$ for Fe(III) and $0.51 \pm 0.31 \mu\text{M}$ for Mn(II), respectively; the range of estimated soluble Fe(III) and Mn(II) was $0.1\sim 10.7 \mu\text{M}$ and $0.05\sim 2.47 \mu\text{M}$. These results are somewhat lower than the values reported by Shao et al. (2019) and other studies cited in the paper but the overall magnitudes are well comparable. We realize that assuming constant Fe and Mn mass fraction is a simplification and latest

CAMx version has the option to treat Fe and Mn as primary species. However, using this option would put even more burden on the emission inventory to have accurate source speciation profiles for different source sectors. Nevertheless, although this Fe(III)/Mn(II) catalyzed pathways stands out among all aqueous pathways under polluted conditions, the relative contribution (8.4%) is only about one third of that from the SO₂ + NO₂ heterogeneous reactions (21.1%). As for the SO₂ + NO₂ heterogeneous reactions, its contribution to sulfate formation doubled from 10.5% (0.21 µg m⁻³) under clean conditions to 21.1% (1.2 µg m⁻³) under polluted conditions. Under all conditions, the relative contribution of the SO₂ + NO₂ heterogeneous reactions exceeds the sum of all aqueous pathways, indicating the importance of heterogeneous oxidation pathways compared to aqueous pathways.

Sulfate formation under selected episodes

We further selected four heavy haze episodes (EP1-EP4) with observed sulfate concentrations continuously exceeding 30 µg/m³ (as highlighted in Figure 4) at the SAES site. These episodes lasted from 9 hours (EP2) to as long as 37 hours (EP1) with episode average sulfate concentrations are all above 50 µg m⁻³ (Figure S7) except for EP3 (36.2 µg m⁻³) (Table S6). Maximum hourly sulfate concentrations ranged from 48.6 µg m⁻³ for EP3 to 93.4 µg m⁻³ for EP2. The averaged molar sulfate and SO₂ ratio ([SO₄²⁻]/[SO₂]) for EP1 and EP2 are higher (0.52 and 0.70, respectively) than that for EP3 (0.17) and EP4 (0.19). In the base case scenario, sulfate formation was significantly underestimated for all four episodes with NMB ranging from -39% to -72%. Figure S8 shows the sulfate formation budget for the four episodes of the base case scenario. The gas phase oxidation pathway was the dominant contributor, accounting for 52% (EP2) to 79% (EP3) of total secondary sulfate formation, followed by the SO₂ + NO₂ heterogeneous reactions with contributions of 20% ~ 39%. For EP1 and EP2, the Fe/Mn-catalyzed oxidation pathway contributed ~10% of sulfate formation but were negligible for the other two episodes. It is interesting to note that for all selected episodes except EP3, sulfate formation was enhanced in scenario Het_2NH₃ by 10.4 to 14.6 µg m⁻³ while EP3 only exhibits minimal increase of modeled sulfate concentrations by only 0.8 µg m⁻³. We performed additional sensitivity tests and found that even with 10 times ammonia emissions, modeled sulfate concentration during EP3 is enhanced by only 2.3 µg m⁻³, which is still much lower compared to the observed values. We suspect that other factors, for example, meteorology might be biased during EP3 and lead to the underpredicted sulfate concentrations. For instance, we looked at the model performance of WRF predictions for individual episode. All four episodes had some over-prediction of wind speeds with NMB ranging from 4% of EP2 to as much as 43% of EP3. Clearly, the large over-prediction of wind speeds during EP3 contributed partially to the underestimated sulfate concentrations by the model. Another potential cause for sulfate underprediction could be failure to capture episodic primary sulfate emissions during EP3. When EP3 is excluded, modeled sulfate concentrations during heavy pollution episodes are greatly enhanced from 33.5 µg m⁻³ in the base scenario to 46.2 µg m⁻³ in scenario Het_2NH₃ (increase by 38 %), due to the combined influences of the SO₂ + NO₂ heterogeneous reactions and doubled ammonia emissions.

3.3 Observed and predicted aerosol pH at the SAES site

Aerosol pH, which is calculated from ISORROPIA either based on observations or within CAMx, is crucial for the heterogeneous $\text{SO}_2 + \text{NO}_2$ reactions to be effective. Observation-based aerosol pH was calculated using forward metastable mode by ISORROPIA to be consistent with CAMx ISORROPIA configuration. Figure 8 shows the distribution of observation-based and modeled aerosol pH at the SAES site by scenario/period. As indicated by both observation-based and modeled pH values, aerosols become more acidic as pollution develops. This is consistent with the higher SO_2 concentrations observed under polluted conditions (Figure S9). For observation-based values, aerosol pH dropped by 35% from clean to polluted conditions while modeled aerosol pH dropped by 13~17% under different scenarios. As also shown by Figure 8, observation-based aerosol pH values are consistently higher than modeled values for all scenarios. Averaged observed-based pH value during clean, transition, and polluted period is 5.5, 4.7, and 3.6, while corresponding value for base scenario (noHet) is 2.8, 2.6 and 2.3, each representing an underestimation by 48%, 45% and 34%. Maximum aerosol pH reached 5.0, 4.4, and 3.8 under clean, transition and polluted periods in the base scenario in contrast to observation-based values of 7.7, 6.5, and 5.3. Adding the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions causes small decrease (0.03~0.07) in predicted aerosol pH. The discrepancies between observation-based and model-based aerosol pH values might be due to significant underprediction of NH_3 and ammonium concentrations. Therefore, when NH_3 emissions are doubled, modeled aerosol pH increases by ~0.7 to 3.0–3.5 and underestimation of aerosol pH for scenario noHet_2 NH_3 is reduced to 36% during clean periods and 15% during polluted periods. Maximum aerosol pH during clean, transition and polluted periods is 5.7, 5.1, and 4.2 under scenario noHet_2 NH_3 . Again, adding the $\text{SO}_2 + \text{NO}_2$ reactions on top of doubled NH_3 emissions slightly decreases the aerosol pH by 0.03–0.12, with stronger reduction associated with more enhancement of sulfate formation. Both observation-based and model-based aerosol pH values at the SAES site indicate that aerosol pH is acidic, which is lower than the more neutralized values reported in previous studies for the Beijing-Tianjin-Hebei region (e.g. 5.4 to 6.2 reported by Cheng et al. (2016) and 6.0 to 7.6 by Wang et al. (2016), the latter was later found to be associated with a coding bug in ISORROPIA). This difference might be due to lower ammonia levels in Shanghai compared with Beijing (S. Wang et al. 2015). However, even when ammonia emissions are increased by 10 times, maximum modeled aerosol pH value is 4.8 under polluted condition, which is still lower than the values reported for north China. Our results indicated that the aerosol pH at the SAES site tends to be moderately acidic regardless of the ambient ammonia concentrations. However, the acidity of aerosols in China still remains to be a vigorous debate. For example, Shi et al. (2017) reported a wide range of pH values between 0.33 and 13.6, depending on the source contributions. Xie et al. (2019) found that the predicted particulate pH values increased from moderate acid to near neutral with the increase of nitrate to sulfate molar ratio.

3.4 Spatial impact of the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions and ammonia emissions

Figure 9 shows the spatial distribution of monthly mean sulfate, SO_2 , ammonia concentrations, and aerosol pH simulated in the base case and the differences between base case and other three sensitivity runs in the YRD region. Similar plots of

nitrate, ammonium, and $\text{PM}_{2.5}$ are shown in Figure S10. Overall, impacts of the heterogeneous reactions and ammonia emissions over the YRD region are similar to that observed at the SAES site. With the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions only, predicted monthly mean sulfate concentrations show ubiquitous increase of $0.1\text{--}5\ \mu\text{g m}^{-3}$ across the domain with larger increase observed in the north and northwest directions. Regions with relative higher increase of predicted sulfate concentrations closely track regions with relatively high aerosol pH and high ammonia concentrations. Aerosol pH decreases slightly because more SO_2 is pulled into the aerosol phase. For nitrate concentrations (Figure S10), however, the heterogeneous reactions lead to increase in the northwest region but decrease for the rest of the YRD region and magnitudes of changes in both directions are within $1\ \mu\text{g m}^{-3}$. Predicted ammonium concentrations show less than $1\ \mu\text{g m}^{-3}$ increase over the majority of the domain. Domain average $\text{PM}_{2.5}$ concentrations increased by $1.2\ \mu\text{g m}^{-3}$ with spatial patterns similar to sulfate.

With doubled ammonia emissions, predictions of all three inorganic PM species are enhanced with most profound impacts observed for nitrate (Figure 8 and Figure S10). Uniform increase across the YRD region is observed for predicted sulfate concentrations; for nitrate and ammonium, increase of predicted concentrations is more significant towards the south. Domain averaged sulfate, nitrate, ammonium and $\text{PM}_{2.5}$ concentrations increase by 0.5, 6.2, 0.3, and $8.0\ \mu\text{g m}^{-3}$, respectively. Aerosol pH is also elevated (on average by 0.3) with more ammonia available. In south Anhui and south Zhejiang provinces, elevation of aerosol pH exceeds one unit. Areas with larger pH increase are also areas with relatively lower pH values in the base scenario, indicating that aerosol pH responds nonlinearly to changes in ammonia emissions.

When both the heterogeneous reactions and doubled ammonia emissions are considered, simulated sulfate concentrations are enhanced by $2.7\ \mu\text{g m}^{-3}$ across the YRD region. Again, areas with relatively larger enhancement of sulfate concentrations are regions with relatively high aerosol pH values and not necessarily regions with maximum increase of aerosol pH. Minimal changes in nitrate and ammonium concentrations are observed with and without the heterogeneous reactions when ammonia emissions are doubled. For $\text{PM}_{2.5}$, domain average concentrations increase by $11.6\ \mu\text{g m}^{-3}$. Simulated $\text{PM}_{2.5}$ concentrations show better agreement with observations at the 24 monitoring sites (Figure 5); averaged NMB is reduced from -21 % in the base scenario to -11 % in scenario Het_2 NH_3 .

Figure 10 further compares the average simulated sulfate concentrations between the base case and Het scenario for the outer 36 km domain during the modelling period. In the base case simulation, high sulfate concentrations were noticed at scattered cities over the North China Plain, Central China and the central part of the Sichuan Basin, corresponding to regions with elevated SO_2 concentrations. Implementing the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions enhanced simulated sulfate concentrations by at least $1\text{--}5\ \mu\text{g m}^{-3}$ for regions to the east of the “Hu Line”. In particular, Northeast China Plain shows most significant sulfate enhancement of more than $10\ \mu\text{g m}^{-3}$; simulated average sulfate concentrations in the Northeast China Plain increased from less than $20\ \mu\text{g m}^{-3}$ during the base case scenario to more than exceed $30\ \mu\text{g m}^{-3}$ in the Het scenario. For other regions, including the North China Plain and Sichuan Basin that show relative high sulfate concentrations in the base case scenario, sulfate concentrations were increased by $5\text{--}10\ \mu\text{g m}^{-3}$ due to the implementation of the reactive SO_2 uptake mechanism. The spatial pattern of sulfate enhancement generally follows that of predicted ammonia concentrations,

once again suggesting the important role of ammonia emissions for this mechanism. Future studies and local sulfate observations are needed to evaluate this mechanism for other parts of China, especially for Northeast China Plain.

4 Conclusions and recommendations

A new parameterization of the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions based on observations in Beijing to improve model simulation of sulfate formation under heavy haze conditions in the YRD region was implemented in this study. Unlike previous studies that only considered the influence of relative humidity on sulfate formation, we also included the impact of aerosol pH in our parameterization. Four CAMx sensitivity runs were conducted to evaluate the importance of the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions as well as ammonia emissions on simulated sulfate concentrations in the YRD region. Base case simulation showed reasonable model performance of sulfate with an overall MB of $-2.7 \mu\text{g m}^{-3}$ but significantly underpredicted sulfate concentrations by 36 % during polluted conditions. Implementation of the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions alone showed slight improvement of sulfate simulation (increase by 6.5 %) under polluted conditions due to acidic aerosol pH. Ammonia concentrations were significantly underestimated by the model. Doubling ammonia emissions alone exhibited a similar impact (sulfate increase by 5.6 %) with that of the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions alone. Nevertheless, aerosol pH increased by 0.7 with doubled ammonia emissions, which enabled the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions to become effective. Thus, in a fourth scenario where both the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions and doubled ammonia emissions were considered, simulated sulfate concentrations during polluted periods increased from $23.1 \mu\text{g m}^{-3}$ in the base case to $29.1 \mu\text{g m}^{-3}$, representing an increase by 26 %. Results for sulfate simulations over entire China shows that for some parts of China, especially the Northeast China Plain, implementing the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions could lead to as much as $20 \mu\text{g m}^{-3}$ increase of sulfate concentrations and the spatial pattern of sulfate enhancement follows closely to that of ammonia concentrations.

In this study, we looked at two questions: (a) is the observation-based estimate from Beijing compatible with data from Shanghai? The answer is yes. Based on the observed data of our modelling period in Shanghai, heavy winter haze events share many similarities between Beijing and Shanghai, for example, increased sulfate fraction under polluted conditions, sulfate to SO_2 ratio increases with RH. If the parameterization based on Beijing data generates too much sulfate in Shanghai, we could conclude that the parameterization is inconsistent with the data for Shanghai. However, this was not the case, even with increase ammonia emissions. So the parameterization is consistent with the data for Shanghai. The second question we wanted to answer is (b) when we apply the rate estimate for Beijing to Shanghai, do we find that $\text{NO}_2 + \text{SO}_2$ can be potentially important in Shanghai? The answer is also yes, depending on the pH level, which depends on the NH_3 emission. However, under-prediction of sulfate concentration still exists (by 20 %) in the YRD region under polluted conditions even with the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions and doubled ammonia emissions, which urges further efforts to better constrain the parameterization of the $\text{SO}_2 + \text{NO}_2$ heterogeneous reactions using local data and to improve the accuracy of ammonia emissions inventory. For instance, we would ideally like to have observed sulfate and precursor concentrations at a pair of

sites oriented upwind and downwind (so that the differences between the two sites constrain sulfate production), observed particle characterization (e.g. particle diameter, number concentrations), and robust NH_3 concentrations to constrain aerosol pH. Follow up studies can be conducted once these data become available. The analysis of the modelled sulfate formation budget shows that transition metal (Fe and Mn) concentrations are influential on sulfate production and therefore improving the emission inventory for $\text{PM}_{2.5}$ Fe and Mn will improve model performance for sulfate and could also influence model assessments of how sulfate concentrations respond to emission management strategies that reduce $\text{PM}_{2.5}$ Fe and Mn concentrations in the YRD region.

Date and code availability. All data and modified CAMx code is available upon request from the corresponding authors.

Competing interest. The authors declare that they have no conflict of interest.

Special issue statement. This article is part of the special issue “Multiphase chemistry of secondary aerosol formation under severe haze”. It is not associated with a conference.

Author contribution. L.H., J.A., L.L., C.H. and G.Y. designed the research; B.K. and L.H. modified the code; R.Y. conducted WRF simulation; J.A. conducted CAMx simulations; L.H. and J.A. analyzed the data; L.L., G.Y., C.H. and Y.W. provided important academic guidance; L.H. and J.A. wrote the paper with contributions from all authors.

Acknowledgement. This study was financially sponsored by the Shanghai Sail Program (NO. 19YF1415600), the National Natural Science Foundation of China (NO. 41875161), and Chinese National Key Technology R&D Program (NO. 2014BAC22B03 and NO. 2018YFC0213800). We thank Qi Zhang, Qian Wang, and Hongli Li from Shanghai University for helping with the data analysis.

References

- Cheng, Y., Zheng, G., Wei, C., Mu, Q., Zheng, B., Wang, Z., Gao, M., Zhang, Q., He, K., Carmichael, G., Pöschl, U., and Su, H.: Reactive nitrogen chemistry in aerosol water as a source of sulfate during haze events in China, *Science Advances.*, 2, e1601530-e1601530, doi:10.1126/sciadv.1601530, 2016.
- Emmons, L. K., Walters, S., Hess, P. G., J.-F., L., Pfister, G. G., Fillmore, D., Granier, C., Guenther, A., Kinnison, D., Laepple, T., Orlando, J., Tie, X., Tyndall, G., Wiedinmyer, C., Baughcum, S. L., and Kloster, S.: Description and evaluation of the model for ozone and related chemical tracers, version 4 (mozart-4), *Geoscientific Model Development.*, 3, 43-67, doi:10.5194/gmd-3-43-2010, 2010.

- Fu, X., Wang, S., Zhao, B., Xing, J., Cheng, Z., Liu, H., and Hao, J.: Emission inventory of primary pollutants and chemical speciation in 2010 for the Yangtze River Delta region, China, *Atmos. Environ.*, 70, 39-50, doi:39-50.10.1016/j.atmosenv.2012.12.034, 2013.
- Gao, M., Carmichael, G. R., Wang, Y., Ji, D., Liu, Z., and Wang, Z.: Improving simulations of sulfate aerosols during winter
5 haze over Northern China: the impacts of heterogeneous oxidation by NO₂, *Frontiers of Environmental Science & Engineering.*, 10, 16, doi:10.1007/s11783-016-0878-2, 2016.
- Gao, M., Carmichael, G. R., Wang, Y., Saide, P. E., Yu, M., Xin, J., Liu, Z., and Wang, Z.: Modeling study of the 2010 regional haze event in the North China Plain, *Atmos. Chem. Phys.*, 16, 1673-1691,doi:10.5194/acpd-15-22781-2015, 2016.
- Guenther, A. B., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T., Emmons, L. K., and Wang, X.: The Model of
10 Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1): an extended and updated framework for modeling biogenic emissions, *Geoscientific Model Development.*, 5, 1471-1492, doi:10.5194/gmd-5-1471-2012, 2012.
- Guo, H., Weber, R. J., and Nenes, A.: High levels of ammonia do not raise fine particle pH sufficiently to yield nitrogen oxide-dominated sulfate production, *Scientific Reports.*, 7, 12109, doi:10.1038/s41598-017-11704-0, 2017.
- He, H., Wang, Y., Ma, Q., Ma, J., Chu, B., Ji, D., Tang, G., Liu, C., Zhang, H., and Hao, J.: Mineral dust and NO_x promote
15 the conversion of SO₂ to sulfate in heavy pollution days, *Scientific Reports.*, 4, 4172, doi:10.1038/srep04172, 2014.
- Huang, Q., Cheng, S., Li, J., Chen, D., Wang, H., and Guo, X.: Assessment of PM₁₀ emission sources for priority regulation in urban air quality management using a new coupled MM5-CAMx-PSAT modeling approach, *Environmental Engineering Science.*, 29, 343-349, doi:http://doi.org/10.1089/ees.2011.0229, 2012.
- Hung, H. M., Hsu, M. N., and Hoffmann, M. R.: Quantification of SO₂ Oxidation on Interfacial Surfaces of Acidic Micro-
20 Droplets: Implication for Ambient Sulfate Formation, *Environmental Science & Technology.*, 52, 9079-9086, doi:10.1021/acs.est.8b01391, 2018.
- Jacobson, M. Z.: Development and application of a new air pollution modeling system—II. Aerosol module structure and design. *Atmospheric Environment*, 31(2), 131-144, 1997.
- Jia, J., Cheng, S., Liu, L., Lang, J., Wang, G., Chen, G., and Liu, X.: An Integrated WRF-CAMx Modeling Approach for
25 Impact Analysis of Implementing the Emergency PM_{2.5} Control Measures during Red Alerts in Beijing in December 2015, *Aerosol and Air Quality Research.*, 17, 2491-2508, doi:10.4209/aaqr.2017.01.0009, 2017.
- Li, L., Chen, C. H., Fu, J. S., Huang, C., Streets, D. G., Huang, H. Y., Zhang, G. F., Wang, Y. J., Jang, C. J., Wang, H. L., Chen, Y. R., and Fu, J. M.: Air quality and emissions in the Yangtze River Delta, China, *Atmos. Chem. Phys.*, 11, 1621-1639, doi:10.5194/acpd-10-23657-2010, 2011.
- 30 Li, L., An, J. Y., Zhou, M., Yan, R. S., Huang, C., Lu, Q., Lin, L., Wang, Y. J., Tao, S. K., Qiao, L. P., Zhu, S. H., and Chen, C. H.: Source apportionment of fine particles and its chemical components over the Yangtze River Delta, China during a heavy haze pollution episode, *Atmos. Environ.*, 123, 415-429, doi:10.1016/j.atmosenv.2015.06.051, 2015.

- Li, L., Cheng, S., Li, J., Lang, J., and Chen, D.: Application of MM5-CAMx-PSAT Modeling Approach for Investigating Emission Source Contribution to Atmospheric Pollution in Tangshan, Northern China, *Mathematical Problems in Engineering.*, 2013, 1-12, doi:10.1155/2013/136453, 2013.
- Liu, Y., Li, L., An, J., Huang, L., Yan, R., Huang, C., Wang, H., Wang, Q., Wang, M., and Zhang, W.: Estimation of biogenic VOC emissions and its impact on ozone formation over the Yangtze River Delta region, China, *Atmos. Environ.*, 186, 113-128, doi:10.1016/j.atmosenv.2018.05.027, 2018.
- Martin, L. R., & Good, T. W.: Catalyzed oxidation of sulfur dioxide in solution: The iron-manganese synergism. *Atmospheric Environment. Part A. General Topics*, 25(10), 2395-2399, 1991.
- Massling, A., Stock, M., Wehner, B., Wu, Z.J., Hu, M., Brüggemann, E., Gnauk, T., Herrmann, H., Wiedensohler, A.: Size segregated water uptake of the urban submicrometer aerosol in Beijing. *Atmospheric Environment* 43 (8), 1578e1589, 2009.
- Ming, L., Ling, J., Li, J., Fu, P., Yang, W., Di, L., Gan, Z., Wang, Z., and Li, X.: PM_{2.5} in the Yangtze River Delta, China: Chemical compositions, seasonal variations, and regional pollution events, *Environ. Pollut.*, 223, 200-212, doi:10.1016/j.envpol.2017.01.013, 2017.
- Quan, J., Tie, X., Zhang, Q., Liu, Q., Li, X., Gao, Y., and Zhao, D.: Characteristics of heavy aerosol pollution during the 2012-2013 winter in Beijing, China, *Atmos. Environ.*, 88, 83-89, doi:10.1016/j.atmosenv.2014.01.058, 2014.
- Ramboll Environ.: User's Guide for Comprehensive Air Quality Model with Extensions Version 6.40. Ramboll Environ, Novato, California, http://www.camx.com/files/camxusersguide_v6-40.pdf, 2016.
- Seinfeld, J. H., & Pandis, S. N.: *Atmospheric chemistry and physics: from air pollution to climate change*. Atmospheric chemistry and physics: from air pollution to climate change, 2006.
- Shao, J., Chen, Q., Wang, Y., Lu, X., He, P., Sun, Y., ... & Zhao, Y.: Heterogeneous sulfate aerosol formation mechanisms during wintertime Chinese haze events: air quality model assessment using observations of sulfate oxygen isotopes in Beijing. *Atmospheric Chemistry and Physics*, 19(9), 6107-6123, 2019.
- Shi, G., Xu, J., Peng, X., Xiao, Z., Chen, K., Tian, Y., ... & Russell, A. G.: pH of aerosols in a polluted atmosphere: source contributions to highly acidic aerosol. *Environmental Science & Technology*, 51(8), 4289-4296, 2017.
- Song, S., Gao, M., Xu, W., Shao, J., Shi, G., Wang, S., ... & McElroy, M. B.: Fine-particle pH for Beijing winter haze as inferred from different thermodynamic equilibrium models. *Atmospheric Chemistry and Physics*, 18(10), 7423-7438, 2018a.
- Song, S., Gao, M., Xu, W., Sun, Y., Worsnop, D. R., Jayne, J. T., Zhang, Y., Zhu, L., Li, M., Zhou, Z., Cheng, C., Lv, Y., Wang, Y., Peng, W., Xu, X., Lin, N., Wang, Y., Wang, S., Munger, J. W., Jacob, D., and McElroy, M. B.: Possible heterogeneous hydroxymethanesulfonate (HMS) chemistry in northern China winter haze and implications for rapid sulfate formation, *Atmos. Chem. Phys. Discuss.*, doi:10.5194/acp-2018-1015, 2018b.
- Wang, G., Zhang, R., Gomez, M. E., Yang, L., Zamora, M. L., Hu, M., Lin, Y., Peng, J., Guo, S., Meng J and Li, J.: Persistent sulfate formation from London Fog to Chinese haze, *Proceedings of the National Academy of Sciences.*, 113, 13630-13635, doi:10.1073/pnas.1616540113, 2016.

- Wang, M., Cao, C., Li, G., and Singh, R. P.: Analysis of a severe prolonged regional haze episode in the Yangtze River Delta, China, *Atmos. Environ.*, 102, 112-121, doi:10.1016/j.atmosenv.2014.11.038, 2015.
- Wang, S., Nan, J., Shi, C., Fu, Q., Gao, S., Wang, D., Cui, H., Saiz-Lopez.,and Zhou, B.: Atmospheric ammonia and its impacts on regional air quality over the megacity of Shanghai, China, *Scientific reports.*, 5, 15842, doi:10.1038/srep15842, 2015.
- Wang, Y., Zhang, Q., Jiang, J., Zhou, W., Wang, B., He, K., Duan, F.,Zhang, Q., Philip, S.,and Xie, Y.: Enhanced sulfate formation during China's severe winter haze episode in January 2013 missing from current models, *Journal of Geophysical Research: Atmospheres.*, 119, 10425-10440, doi:10.1002/2013jd021426, 2014.
- Wang, X., Li, J., Zhang, Y., Xie, S., and Tang, X.: Ozone source attribution during a severe photochemical smog episode in Beijing, China, *Science in China Series B: Chemistry*, 52, 1270-1280, doi:10.1007/s11426-009-0137-5, 2009.
- Xie, Y., Wang, G., Wang, X., Chen, J., Chen, Y., Tang, G., Wang, L., Ge, S., Xue, G., Wang, Y., and Gao, J.: Observation of nitrate dominant PM_{2.5} and particle pH elevation in urban Beijing during the winter of 2017, *Atmos. Chem. Phys. Discuss.*, <https://doi.org/10.5194/acp-2019-541>, in review, 2019.
- Xu, J. S., Xu, H. H., Xiao, H., Tong, L., Snape, C. E., Wang, C. J., and He, J.: Aerosol composition and sources during high and low pollution periods in Ningbo, China, *Atmos. Res.*, 178-179, 559-569, doi:10.1016/j.atmosres.2016.05.006, 2016.
- Ye, X.N., Ma, Z., Hu, D.W., Yang, X., Chen, J.M.: Size-resolved hygroscopicity of submicrometer urban aerosols in Shanghai during wintertime. *Atmospheric Research* 99 (2), 353e364, 2011.
- Ye, X., Tang, C., Yin, Z., Chen, J., Ma, Z., Kong, L., ... & Geng, F.: Hygroscopic growth of urban aerosol particles during the 2009 Mirage-Shanghai Campaign. *Atmospheric environment*, 64, 263-269, 2013.
- Yarwood, G., J. Jung, G. Z. Whitten, G. Heo, J. Mellberg and E. Estes.: Updates to the Carbon Bond Mechanism for Version 6 (CB6), Presented at the 9th Annual CMAS Conference, Chapel Hill, October, 2010.
- Zhang, L., Brook, J. R., and Vet, R.: A revised parameterization for gaseous dry deposition in air-quality models, *Atmospheric Chemistry and Physics*, 3, 2067-2082, doi:10.5194/acp-3-2067-2003, 2003.
- Zheng, B., Zhang, Q., Zhang, Y., He, K. B., Wang, K., Zheng, G. J., Duan, F. K., Ma, Y. L., and Kimoto, T.: Heterogeneous chemistry: a mechanism missing in current models to explain secondary inorganic aerosol formation during the January 2013 haze episode in North China, *Atmos. Chem. Phys.*, 15, 2031-2049, doi:10.5194/acp-15-2031-2015, 2015a.
- Zheng, G. J., Duan, F. K., Su, H., Ma, Y. L., Cheng, Y., Zheng, B., Zhang, Q., Huang, T., Kimoto, T., Chang, D., Pöschl, U., Cheng, Y. F., and He, K. B.: Exploring the severe winter haze in Beijing: the impact of synoptic weather, regional transport and heterogeneous reactions, *Atmos. Chem. Phys.*, 15, 2969-2983, 10.5194/acp-15-2969-2015, 2015b.

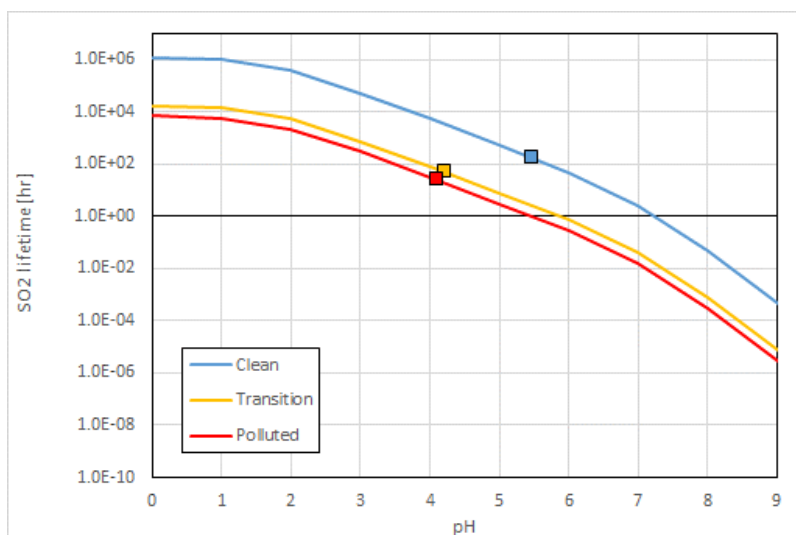
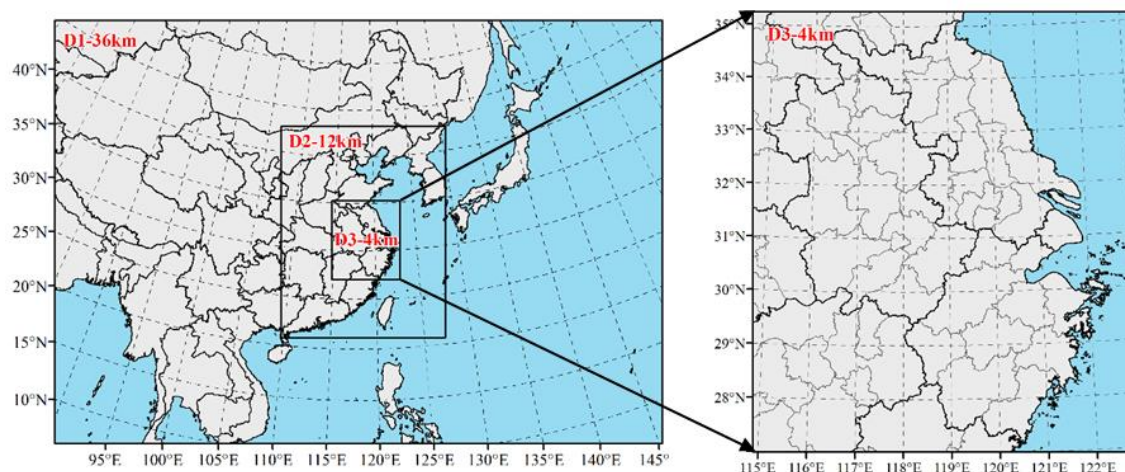


Figure 1: SO_2 lifetime (in hr^{-1}) due to $\text{SO}_2 + \text{NO}_2$ reactive uptake mechanism as a function of aerosol pH under clean, transition, and polluted conditions. Values of relative humidity, temperature, and $\text{NO}_2(\text{g})$ concentrations are based on values in Table S1.



5 Figure 2: CAMx model domains

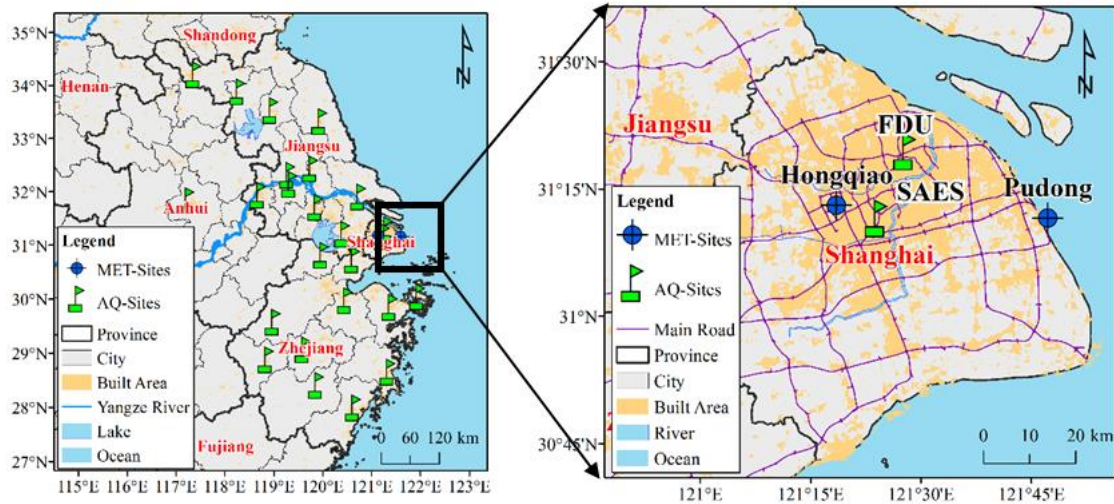


Figure 3: Locations of observations sites for WRF (two MET-Sites) and CAMx model performance evaluation (SAES site and FDU site within Shanghai; another 23 AQ-sites distributed over Jiangsu, Zhejiang, and Anhui province with locations shown in Table S2).

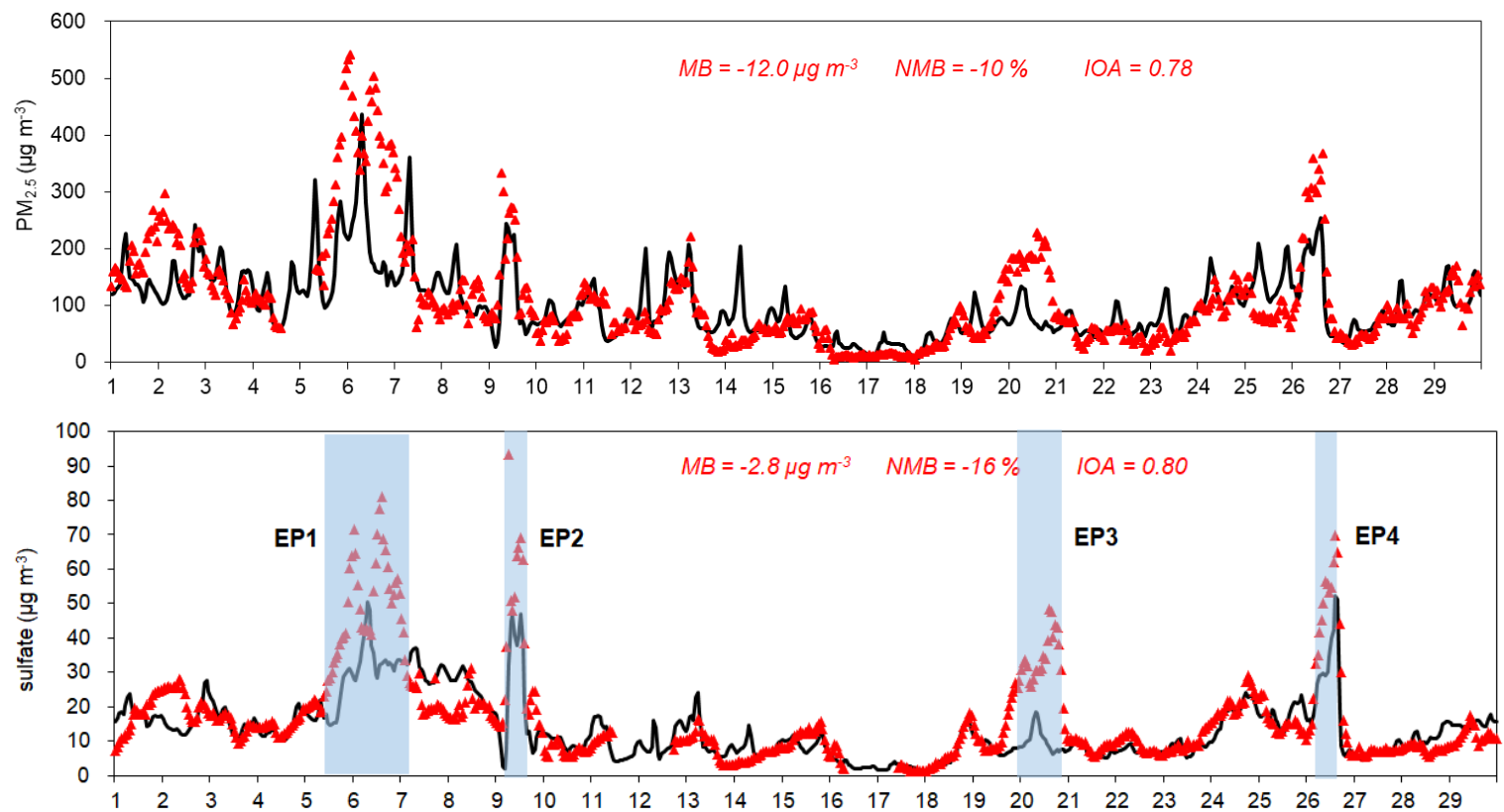


Figure 4: Simulated and observed PM_{2.5} (upper) and sulfate (bottom) concentrations (in µg m⁻³) at SAES site during 1 to 29 December 2013

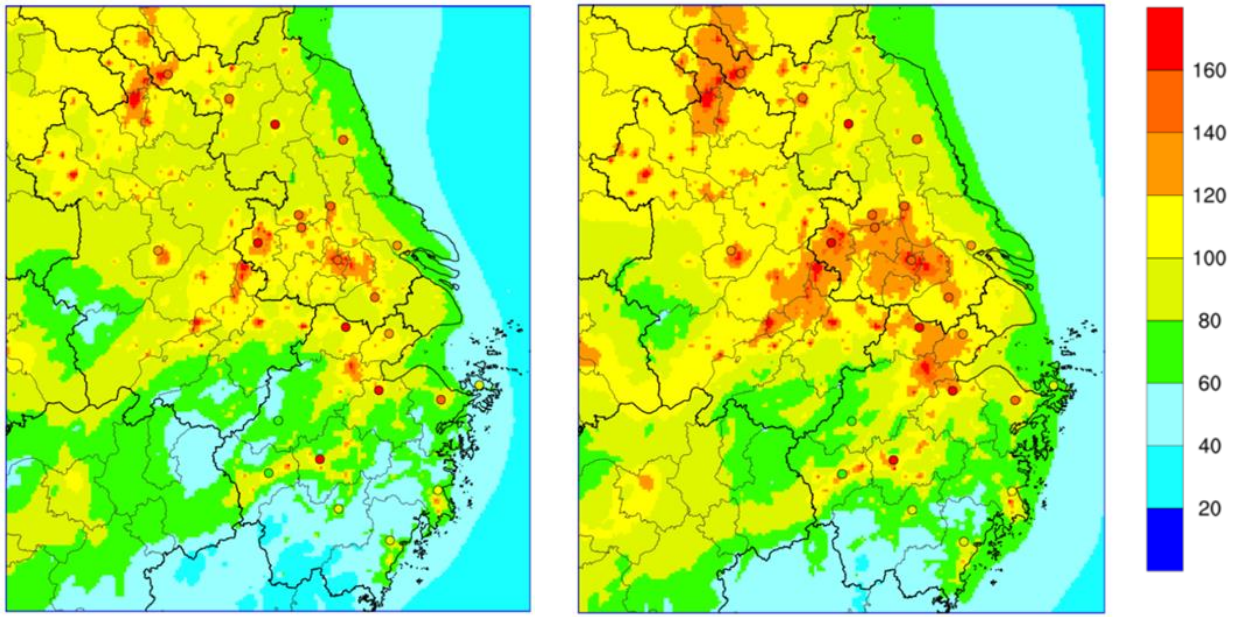


Figure 5: Spatial distribution of observed and simulated monthly average $PM_{2.5}$ concentrations (in $\mu g m^{-3}$) over the YRD region for the base case scenario (left) and Het_2NH₃ scenario (right). Locations of the monitoring sites are listed in Table S2.

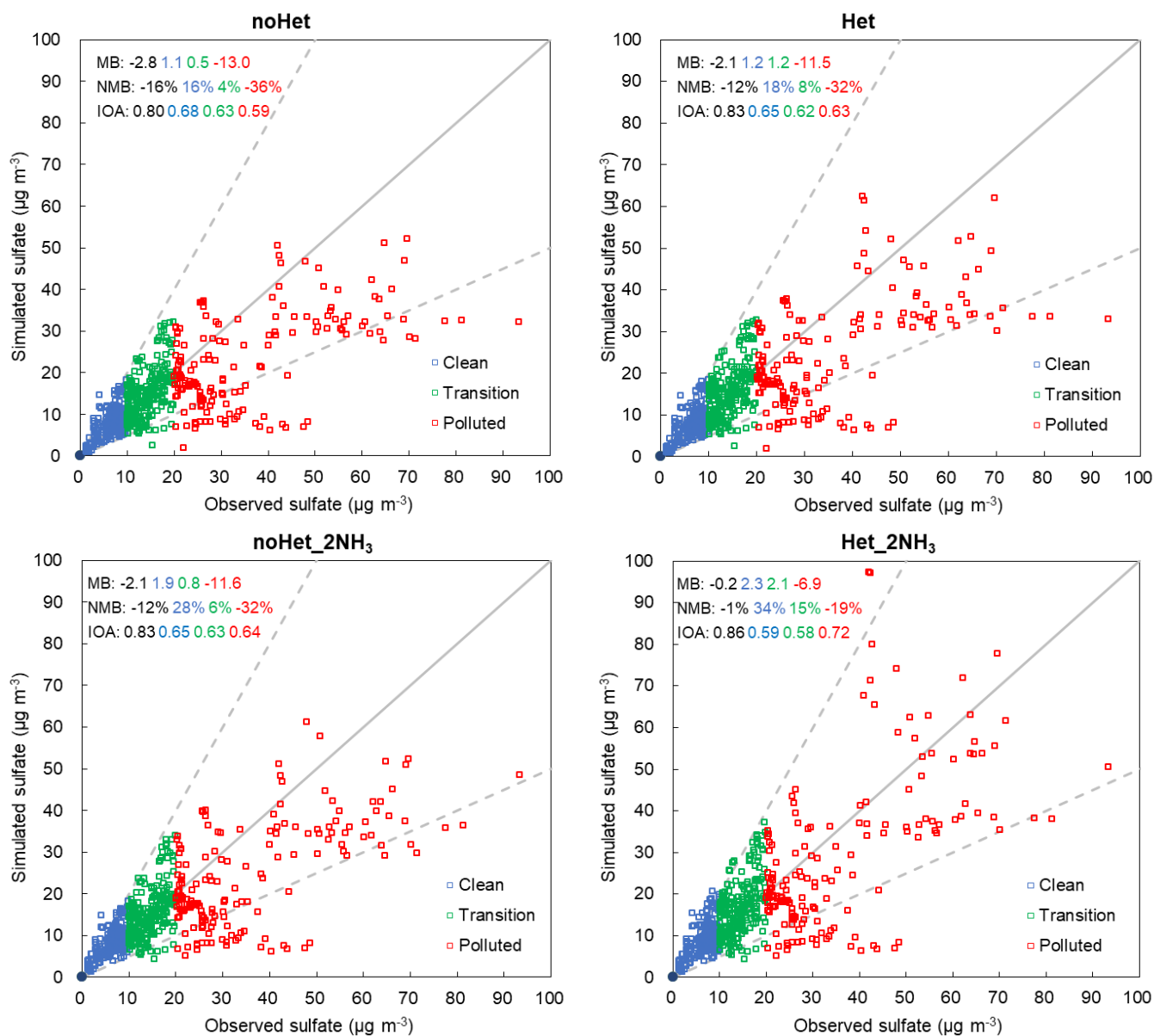


Figure 6: Scatter plots of hourly sulfate concentrations for different scenarios at SAES site during 1 to 29 December 2013. Solid lines indicate 1:1 lines and dashed lines are 1:2 and 2:1 lines.

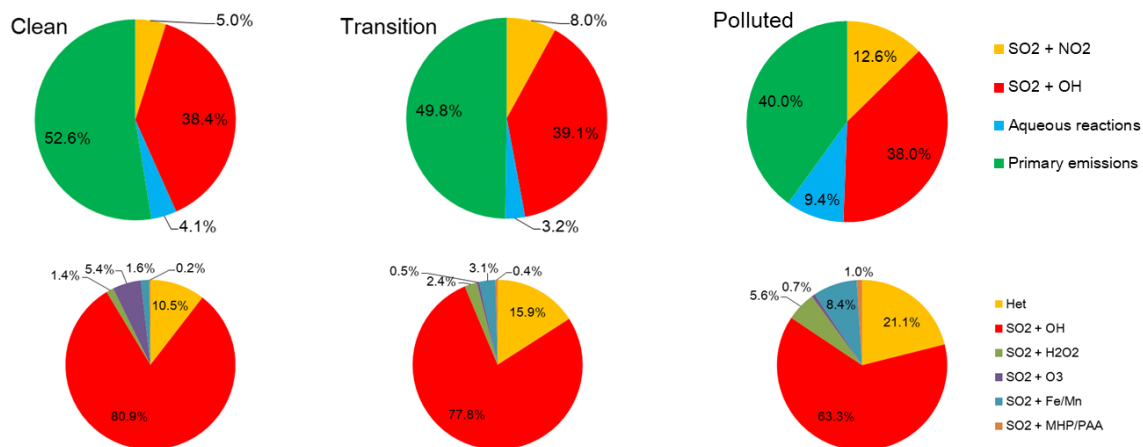
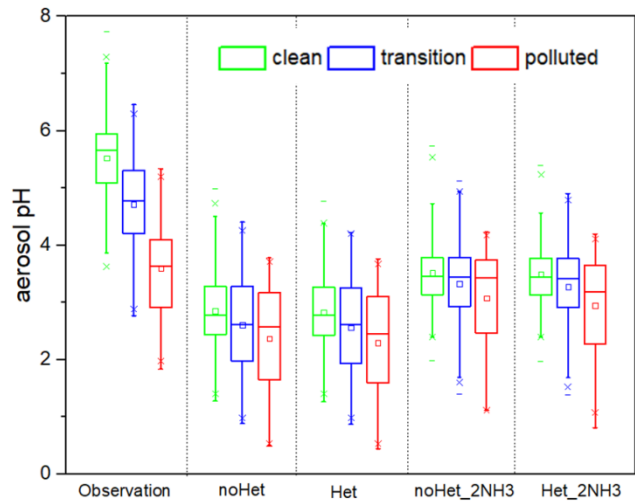


Figure 7: Relative contribution of different pathways to sulfate concentrations at the SAES site during clean, transition and polluted periods. Primary sulfate emissions were excluded in the bottom row.



5 Figure 8: Box and whisker plot of observed and predicted aerosol pH by scenario and period.

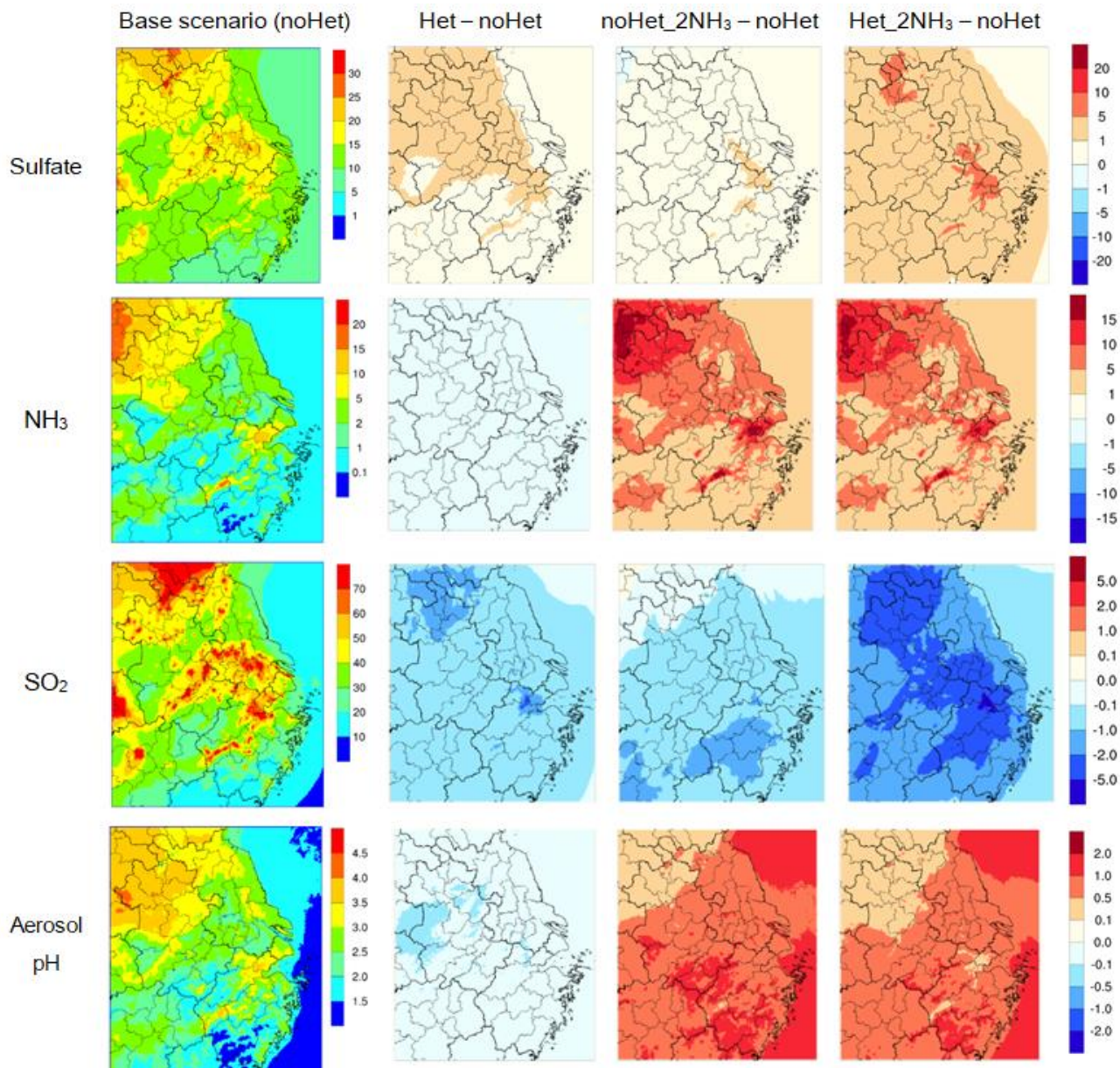


Figure 9: Spatial distribution of simulated monthly sulfate (top row), NH_3 (second row), and SO_2 (third row) concentrations ($\mu\text{g m}^{-3}$) and aerosol pH (bottom row) over the YRD region for the base case scenario (first column) and the differences ($\mu\text{g m}^{-3}$ for concentrations) between the base case and other three scenarios: Het (second column), noHet_2 NH_3 (third column) and Het_2 NH_3 (fourth column).

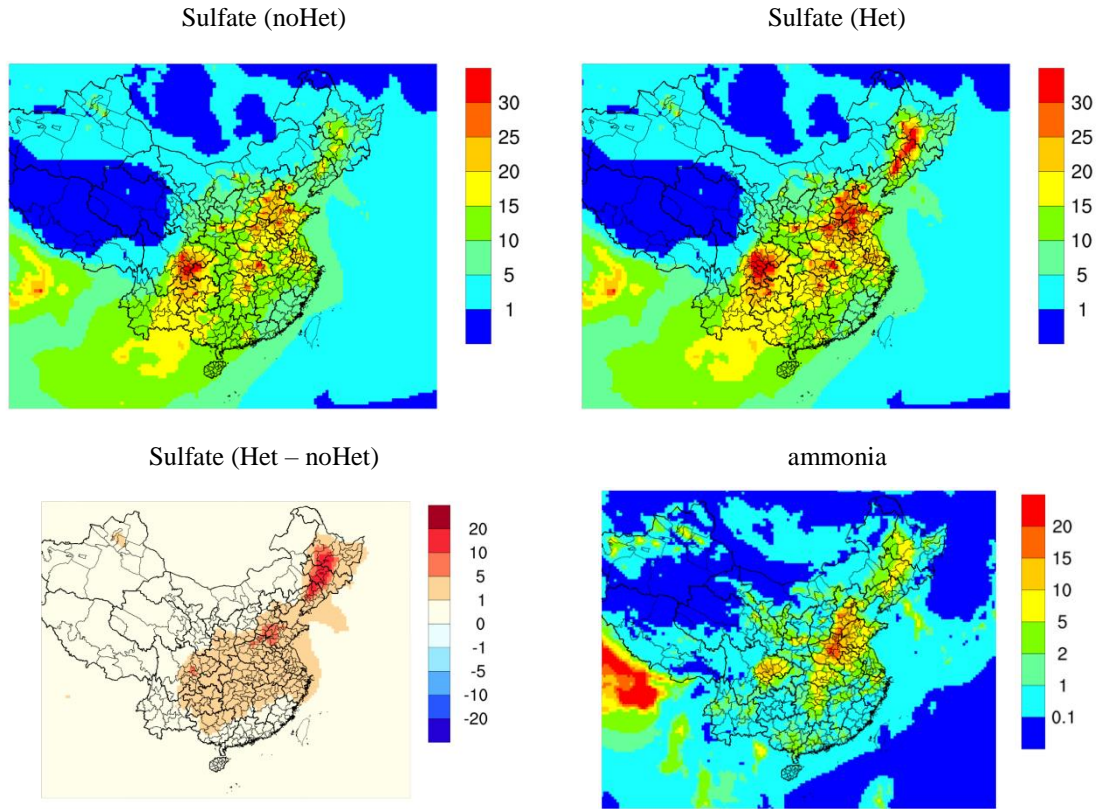


Figure 10: Spatial distribution of simulated monthly sulfate concentrations ($\mu\text{g m}^{-3}$) over the 36 km domain for the base case scenario (top left), Het (top right) and the differences between the two scenarios (bottom left) and ammonia concentrations (in $\mu\text{g m}^{-3}$; bottom right).

5 Table 1. Sulfate formation pathways currently implemented in CAMx version 6.40

No.	Oxidants	Rate expression	Reference
<i>Gaseous Phase</i>			
1	OH	$k_1[\text{OH}][\text{SO}_2(\text{g})]$ $k_1 = \left[\frac{k_0[M]}{1 + \frac{k_0[M]}{k_\infty}} \right] 0.6^G$ $G = \{1 + [\log\left(\frac{k_0[M]}{k_\infty}\right)]^2\}^{-1}$ $k_0 = 4.50 \times 10^{-31} (\text{T}/300)^{-3.9}$ $k_\infty = 1.30 \times 10^{16} (\text{T}/300)^{-0.7}$	Seinfeld and Pandis (2006)
<i>Aqueous Phase</i>			

2	O ₃	$(k_2[\text{SO}_2 \cdot \text{H}_2\text{O}] + k_3[\text{HSO}_3^-] + k_4[\text{SO}_3^{2-}])[\text{O}_3(\text{aq})]$ $k_2 = 2.4 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$ $k_3 = 3.7 \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ $k_4 = 1.5 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$	Jacobson (1997)
3	H ₂ O ₂	$k_5[\text{H}^+][\text{HSO}_3^-][\text{H}_2\text{O}_2(\text{aq})]/(1+13x[\text{H}^+])$ $k_5 = 7.45 \times 10^7 \times \exp(-15.96 \times (298/T - 1)) \text{ M}^{-1} \text{ s}^{-1}$	Jacobson (1997)
4	Fe(III)/Mn(II)	<p>when aerosol pH <=4.2:</p> $750[\text{Mn(II)}][\text{S(IV)}] + 2600[\text{Fe(III)}][\text{S(IV)}] - k_6$ $[\text{Mn(II)}][\text{Fe(III)}][\text{S(IV)}][\text{H}^+]^{0.67}$ <p>when aerosol pH > 4.2:</p> $750[\text{Mn(II)}][\text{S(IV)}] + 2600[\text{Fe(III)}][\text{S(IV)}] -$ $k_7[\text{Mn(II)}][\text{Fe(III)}][\text{S(IV)}][\text{H}^+]^{-0.74}$ $k_6 = 2.51 \times 10^{13} \text{ M}^{-2} \text{ s}^{-1}$ $k_7 = 3.72 \times 10^7 \text{ M}^{-2} \text{ s}^{-1}$	Martin and Good (1991)
5	Methyl hydroperoxide (CH ₃ OOH) and other organic hydroperoxides	$k_8[\text{H}^+][\text{HSO}_3^-][\text{CH}_3\text{OOH}(\text{aq})]$ $k_8 = 1.90 \times 10^7 \times \exp(-12.75 \times (298/T - 1)) \text{ M}^{-2} \text{ s}^{-1}$	Jacobson (1997)
6	Peracetic acid (CH ₃ C(O)OOH) and other organic peracids	$k_9[\text{H}^+][\text{HSO}_3^-][\text{CH}_3\text{C(O)OOH}(\text{aq})]$ $k_9 = 3.60 \times 10^7 \times \exp(-13.42 \times (298/T - 1)) \text{ M}^{-2} \text{ s}^{-1}$	Jacobson (1997)
<i>Aerosol Aqueous Phase (implemented as pseudo gas phase)</i>			
7	NO ₂	$k_{10}[\text{SO}_2(\text{g})]$ $k_{10} = \left(\frac{d_p}{2D} + \frac{4}{v\gamma} \right)^{-1} S_p$ $\gamma_{\text{low}} = 2 \times 10^{-5}$ $\gamma_{\text{high}} = 5 \times 10^{-5}$	Zheng (2015)

Table 2 Observed and simulated sulfate concentrations ($\mu\text{g m}^{-3}$) for different scenarios by clean, transition, polluted periods at SAES site during 1 to 29 December 2013

Period	Observed	noHet	Het	noHet_2NH ₃	Het_2NH ₃
all	17.2	14.4	15.1	15.2	17.0
clean	6.7	7.8	8.0	8.6	9.1
transition	14.3	14.7	15.3	15.0	16.3
polluted	36.1	23.1	24.6	24.5	29.1