# **Supplementary information**

### S1. Methodology for estimation of the mass concentrations of PM<sub>2.5</sub> components

### S1.1 Organic matter

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The mass concentration of organic matter (OM) was calculated from organic carbon (OC) measurements by multiplying OC by a factor that represents the mass contributions of other elements, such as oxygen, hydrogen, and nitrogen. The OM/OC ratio varies from 1.4 to 2.2 and is expected to increase as aerosols age (El-Zanan et al., 2005). We chose a factor of 1.6 to calculate OM in Beijing following advice in the literature (Xing et al., 2013).

#### S1.2 Primary organic carbon

The two main sources of OC are primary (POC) emissions from fuel combustion, biomass burning, or vehicle exhaust, and secondary organic carbon (SOC) from the oxidation of volatile organic compounds (VOCs) and semi-volatile organic compounds (SVOCs). Elemental carbon (EC) is emitted directly into the atmosphere and has been used as a tracer to estimate co-emitted POC. Meanwhile, SOC can be calculated as follows (Saylor et al., 2006; Turpin and Huntzicker, 1995):

$$SOC = OC - POC = OC - \left[ \left( \frac{OC}{EC} \right)_{\text{pri}} EC + b \right],$$
 (Eq. 1)

where  $(OC/EC)_{pri}$  is the primary OC/EC ratio of combustion sources, and b is the primary OC emitted from non-combustion sources. These two parameters can be derived from a linear regression of OC and EC (Fig. S1). In our study, Deming regression (Deming, 1943) was applied to the lowest 10% of OC/EC values for each season to obtain the two parameters, with the slope as  $(OC/EC)_{pri}$  and the intercept as b (Fig. S1). The regression lines for the four seasons are shown in Fig. S1.

### S1.3 Estimates of the OC/EC ratio in primary sources

Deming regression (1943) was applied to the lowest 10% of OC/EC samples for each season. The scatter plots and regression equations for OC against EC in each season are depicted in Fig. S1. The primary OC/EC ratios were determined to be 3.46,

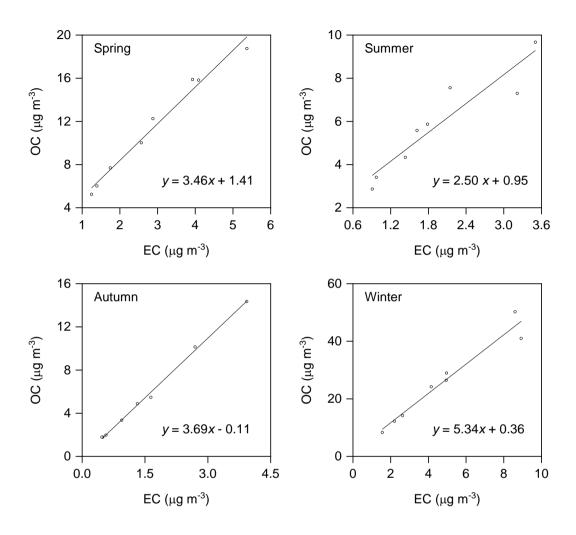


Figure S1. Scatter plots and regression equations for organic carbon (OC) against elemental carbon (EC) in each season.

### S1.4 Minerals

The total mass concentration of minerals, referred to as "minerals", can be estimated by the following equation (Chan et al., 1997):

$$[minerals] = 2.2[Al] + 2.49[Si] + 1.63[Ca] + 2.42[Fe] + 1.94[Ti]$$
, (Eq. 2)

where [x] represents the mass concentration of species x. According to Zhang et al. (2003), on average Al accounted 7% of total mineral dust mass concentrations in North, Northwest, and West China. Mineral concentrations can thus also be estimated by Eq. 3:

$$[minerals] = [A1]/0.07$$
, (Eq. 3)

We calculated [minerals] with the two methods above and found no significant differences (Fig. S2). Equation 3 was therefore employed to calculate [minerals] in this study.

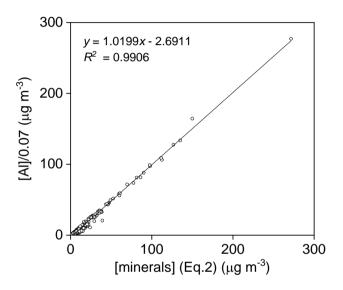


Figure S2. Comparison of the two methods for the calculation of [minerals].

### S1.5 Trace element oxides

The enrichment factors (EFs) of trace element oxides (TEOs) can be used to determine whether natural or anthropogenic sources dominated our observations. The EF value of element *i* was defined as follows:

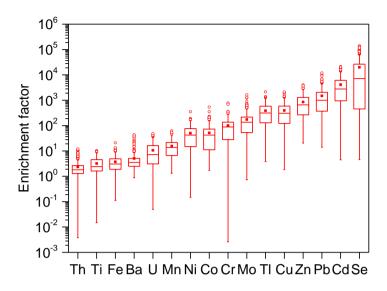
$$EF_i = \frac{[X_i/X_{\text{refl}_{\text{sample}}}}{[X_i/X_{\text{refl}_{\text{crust}}}]},$$
 (Eq. 4)

where  $[X_i/X_{ref}]_{sample}$  is the mass concentration ratio of element i to the reference element in our samples and  $[X_i/X_{ref}]_{crust}$  is the mass concentration ratio of element i to the reference element in average crust (Hans Wedepohl, 1995). All was used as the

reference element in this study. The EFs of each element are depicted in Fig. S3.

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**Figure S3.** Elemental enrichment factors (EFs) of our samples. The boxes represent, from top to bottom, the 75<sup>th</sup>, 50<sup>th</sup>, and 25<sup>th</sup> percentiles for each element. The whiskers, solid red squares, and open red circles represent 1.5 times the interquartile range (IQR), seasonal mean values, and outlier data points, respectively.

If the EF was < 5, the element was considered to originate mainly from natural sources; if 5 < EF < 20, the element originated from both natural and anthropogenic sources; if EF > 20, the element originated mainly from anthropogenic sources. According to Zhang et al. (2013), the mass concentrations of TEOs can be estimated by multiplied a factor to represent the contribution of oxygen. For elements originating from anthropogenic sources only, a factor of 1 was applied, whereas for elements of both natural and anthropogenic origin, a factor of 0.5 was applied to represent the anthropogenic part. As multiple forms of metal oxides were identified, which were hard to quantify, a multiplicative factor of 1.3 was used when considering the metal abundance. The mass concentration of TEOs was calculated as described in Zhang et al. (2013):

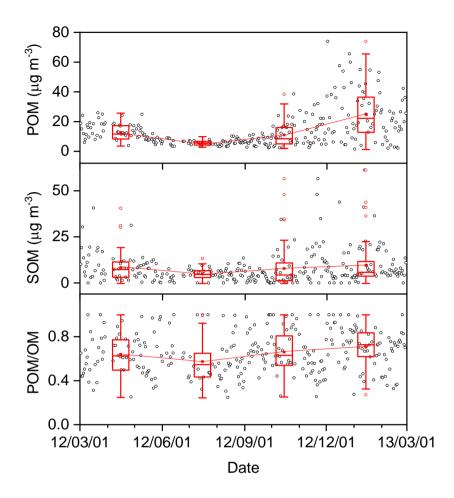
$$[TEOs] = 1.3 \times [0.5 \times (Ba + Mn + U) + (Ni + Co + Cr + Mo + Tl + Cu + Zn + Pb + Cd + Se)],$$
 (Eq. 5)

# S1.6 Aerosol water content

Aerosol water content (AWC) was calculated using the ISORROPIA-II thermodynamic model (http://isorropia.eas.gatech.edu). The  $Na^+-K^+-Ca^{2+}-Mg^{2+}-NH_4^+-SO_4^{2-}-NO_3^--Cl^--H_2O$  aerosol system was applied in reverse mode (Fountoukis and Nenes, 2007; Nenes et al., 1998).

### 5 S2 Results and discussion

### **S2.1** General description



**Figure S4.** Time series of primary organic matter (POM), secondary organic matter (SOM), and POM/OM from March 1 2012 to February 28 2013 (open black circles). The boxes represent, from top to bottom, the 75<sup>th</sup>, 50<sup>th</sup>, and 25<sup>th</sup> percentiles for each season. The whiskers, solid red squares, and open red circles represent 1.5 times the IQR, seasonal mean values, and outlier data points, respectively.

### 5 S2.2 Sulfate formation mechanism

Sulfate can be formed through the oxidation of SO<sub>2</sub> by OH radicals in the gas phase (Stockwell and Calvert, 1983), through the oxidation of dissolved SO<sub>2</sub> by various oxidants (e.g., O<sub>3</sub>, H<sub>2</sub>O<sub>2</sub>, NO<sub>2</sub>, and O<sub>2</sub>) in the aqueous phase (Seinfeld and Pandis, 2006), which may be transition metal ions (TMIs)-catalysed, or through heterogeneous reaction on the surface of sea-salt or

dust aerosols (Gurciullo et al., 1999; Usher, 2002).

The rate of the  $SO_2 + OH$  reaction can be expressed as:

$$R_{SO_2+OH} = k_0[SO_2(g)][OH(g)]$$
, (Eq. 6)

where  $k_0$  is the rate constant and [x] represents the concentration of species x. The production rate of sulfate through OH radical

5 oxidation can be expressed as:

$$P_{\text{OH}} = \frac{3600 \times 96 \times p \times R_{\text{SO}_2 + \text{OH}}}{RT}$$
, (Eq. 7)

where 3600 is a time conversion factor (s h<sup>-1</sup>), 96 is the molar mass of  $SO_4^{2-}$  (g mol<sup>-1</sup>), p is atmospheric pressure (kPa), R is the gas constant (8.31 Pa m<sup>3</sup> mol<sup>-1</sup> K<sup>-1</sup>), and T is the temperature (K).

 $SO_2$  reacts with  $H_2O_2$ ,  $O_3$ ,  $NO_2$ , and  $O_2$  (TMIs-catalysed) in the aqueous phase. The rates of the four main aqueous reactions

are expressed as (He et al., 2018; Seinfeld and Pandis, 2006):

$$R_{SO_2+O_3} = (k_1[SO_2 \cdot H_2O] + k_2[HSO_3^-] + k_3[SO_3^{2-}])[O_3(aq)],$$
 (Eq. 8)

$$R_{\text{SO}_2+\text{H}_2\text{O}_2} = \frac{k_4[\text{H}^+][\text{HSO}_3^-][\text{H}_2\text{O}_2(\text{aq})]}{1+K[\text{H}^+]} , \tag{Eq. 9}$$

$$R_{SO_2+NO_2} = k_5[S(IV)][NO_2(aq)]$$
, (Eq. 10)

$$R_{SO_2+O_2} = k_6 [H^+]^{-0.74} [S(IV)][Mn(II)][Fe(III)]$$
 (pH < 4.2), (Eq. 11)

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$$R_{SO_2+O_2} = k_7 [H^+]^{0.67} [S(IV)][Mn(II)][Fe(III)]$$
 (pH > 4.2), (Eq. 12)

The production rate of sulfate through aqueous oxidation routes can be expressed as:

$$P_{\text{aqu(ox_i)}} = 3600 \times 96 \times R_{\text{SO}_2 + \text{ox}_i} \times \frac{\text{LWC}}{\rho_{\text{H}_2\text{O}}}$$
 (Eq. 13)

where  $k_n$  (n = 1-7) is the rate constant of each oxidation route,  $K = 13 \text{ M}^{-1}$  at 298 K, LWC is the liquid water content (mg m<sup>-3</sup>),  $\rho_{\text{H2O}}$  is the density of water (1 kg L<sup>-1</sup>), and ox<sub>i</sub> ( $i = O_3$ , H<sub>2</sub>O<sub>2</sub>, NO<sub>2</sub>, and O<sub>2</sub>) represents different oxidants.

The heterogeneous reaction rate  $R_{het(ox_i)}$  can be expressed as (Jacob, 2000; Wang et al., 2012; Zheng et al., 2015):

$$R_{\text{het}(ox_i)} = k_{ox_i}[SO_2(g)] , \qquad (Eq. 14)$$

where

$$k_{\text{ox}_i} = \left(\frac{d_{\text{p}}}{2D_{\text{i}}} + \frac{4}{v_i y_i}\right)^{-1} S_{\text{p}} ,$$
 (Eq. 15)

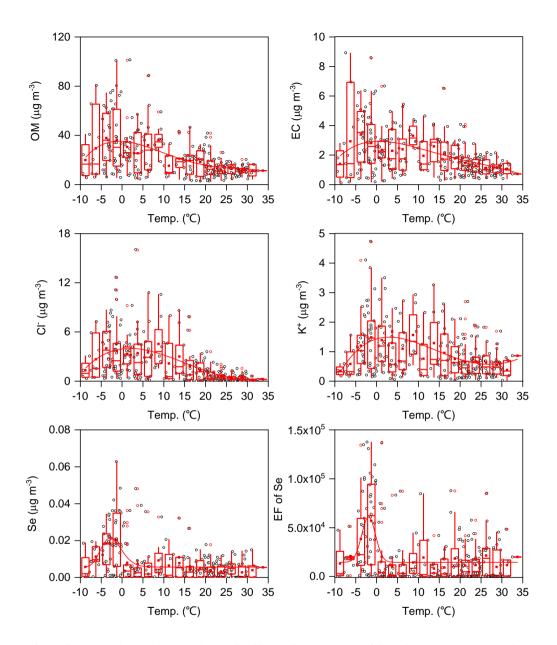
 $d_p$  is the effective diameter of the particles (m),  $D_i$  is the gas phase molecular diffusion coefficient (m<sup>2</sup> s<sup>-1</sup>),  $v_i$  is the mean molecular speed in the gas phase (m s<sup>-1</sup>), and  $S_p$  is the aerosol surface area (m<sup>2</sup> m<sup>-3</sup>). The uptake coefficient  $\gamma_i$  depends on RH:

$$\gamma_{i} = \begin{cases}
\gamma_{\text{low}} & 0 < \text{RH} \le 50\% \\
\gamma_{\text{low}} + \frac{(\gamma_{\text{high}} \cdot \gamma_{\text{low}})(\text{RH} \cdot 0.5)}{\text{RH}_{\text{max}} \cdot 0.5} & 50\% < \text{RH} \le \text{RH}_{\text{max}} \\
\gamma_{\text{high}} & \text{RH}_{\text{max}} < \text{RH} \le 100\%
\end{cases}$$
(Eq. 16)

where  $\gamma_{\text{low}}$  and  $\gamma_{\text{high}}$  can be obtained from Wang et al. (2012) and RH<sub>max</sub> is the RH at which  $\gamma$  reaches  $\gamma_{\text{high}}$ . The rate of sulfate production via heterogeneous reactions  $P_{\text{het}(ox_i)}$  can be expressed as:

$$P_{\text{het}(ox_i)} = \frac{3600 \times 96 \times p \times R_{\text{het}(ox_i)}}{RT} , \qquad (Eq. 17)$$

### **S2.3 Influencing parameters**



**Figure S5.** Plots of organic matter (OM), element carbon (EC),  $Cl^-$ ,  $K^+$ , Se, and the EF of Se against temperature. The boxes represent, from top to bottom, the 75<sup>th</sup>, 50<sup>th</sup>, and 25<sup>th</sup> percentiles in each temperature bin ( $\Delta T = 2.5$  °C). The whiskers, solid red squares, and open red circles represent 1.5 times the IQR, seasonal mean values, and outlier data points, respectively. The red lines are best fits to the mean values based on either polynomial or Gaussian functions.

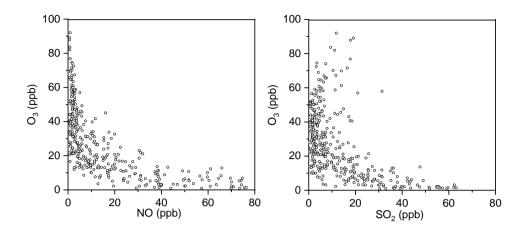


Figure S6. Plots of O<sub>3</sub> against the primary emission tracers NO and SO<sub>2</sub>.

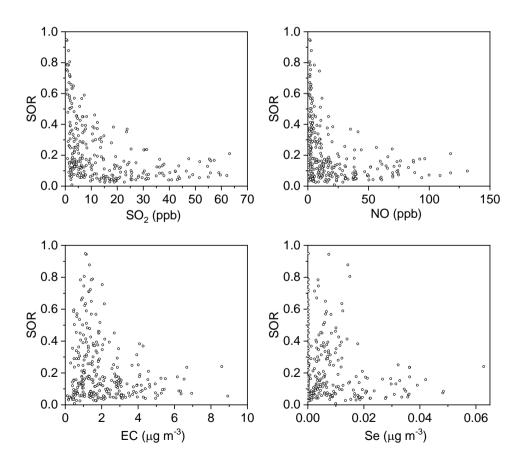
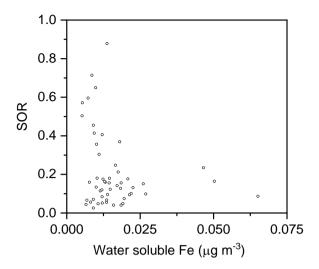
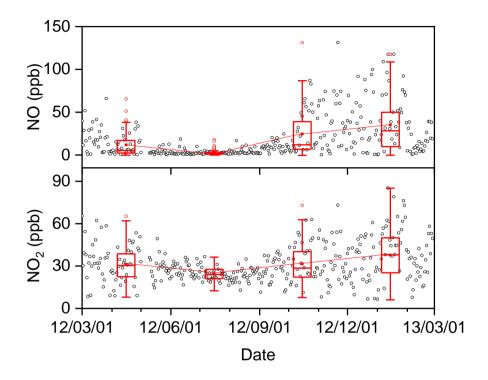


Figure S7. Plots of sulfur oxidation ratios (SORs) against the primary emission tracers SO<sub>2</sub>, NO, EC and Se.



**Figure S8.** Plot of the SOR against water soluble Fe (54 samples selected every 6 days throughout the sampling period).

# **S2.4 Seasonal variations**



**Figure S9.** Time series of NO and NO<sub>2</sub> from March 1 2012 to February 28 2013 (open black circles). The boxes represent, from top to bottom, the 75<sup>th</sup>, 50<sup>th</sup>, and 25<sup>th</sup> percentiles for each season. The whiskers, solid red squares, and open red circles represent 1.5 times the IQR, seasonal mean values, and outlier data points, respectively.

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