

Replies to Referee #1

The authors would like to thank Anonymous Referee #1 for his/her comments.

The Referee's comments (in *italic*) followed by our replies are listed below.

Main concerns: - *Among the main findings of the study is the quantification of biomass burning SOA that is considered to result from aqueous phase reactions of biomass burning products (bb-aqSOA). This consideration, however, is mainly based on high correlations of one of the PMF OOA factors with ALWC. Even though plausible, I hesitate to take this correlation as sufficient evidence for really inferring the formation pathway of this OOA fraction, especially, as no alternative explanations are even discussed by the authors. One alternative explanation would be a strong shift in phase partitioning of water soluble organic gases during periods with high ALWC. A number of compounds like small carboxylic acids and (di)carbonyls present in the gas phase of (aged) BB plumes could be expected to preferentially partition into the particle phase only under high ALWC conditions, owing to their high volatility combined with high water solubility. These would contribute to the "OOAx_BB_aq" even without any aqueous phase chemistry taking place and OOAx_BB_aq might in fact just represent an OOA fraction with different phase partitioning behaviour than OOAx_BB", regardless of their respective formation pathways. There might be other possible explanations for the observed correlation with ALWC and I would recommend providing a more critical discussion on the subject.*

Authors Reply. The authors thank the Reviewer #1 for the critical view and suggestion. We acknowledge that a more clear explanation of our hypothesis can be profitable and it will be added in section 4.3 of the revised manuscript. We modified the manuscript in order to clarify that the attribution of OOAx_BB_aq factor to aqueous phase chemistry is not solely based on the correlation with ALWC.

The consideration that the OOAx_BB-aq factors result dominantly from aqueous-phase reactions is still the most credible for two main reasons:

1- a technical one, the analyzed aerosol was dried to about 35-40% by means of a Nafion drier before sampling with the HR-TOF-AMS. So, most of the possible volatile compounds dissolved into particle phase under high ALWC conditions are expected to be volatilized again during sampling, leaving into aerosol only the most stable (complexed/oxidized by ageing) compounds;
2- linked with the previous, even if it's true that the partitioning of gaseous soluble species emitted by biomass burning is enhanced during high ALWC conditions, it is even true that these small and volatile compounds need to react to become stable components of particle phase. And this is exactly the mechanism that the Authors suggest is happening.

The correlation of OOAx_BB-aq factors with HMSA is a good example of this mechanism, as also the Reviewer #1 suggested in the subsequent comment: "*Formaldehyde would preferentially partition into particles for high ALWC conditions to form HMSA*" and ONLY if it happens the products of this reaction (i.e., HMSA) will remain in the particles after water evaporation. And this is supposed to happen to other gaseous species and their specific products, changing permanently the chemical composition of the organic aerosol.

- *Further evidence provided to corroborate an aqueous phase formation pathway of OOA_BB_aq is either not conflicting with the above given alternative explanation or not fully convincing to me. Formaldehyde would preferentially partition into particles for high ALWC conditions to form HMSA (P13L2), taking AMS fragment ions as evidence for specific compounds in a complex ambient mixture might be questionable (P13L22 and P13L26) and many of these are well correlated to other OOAx factors as well (Table S16), the agreement with fog processing spectra is only partial in many cases (P13L18 and Table S15) and O:C ratios are unlikely to be good indicators for specific formation pathways (P13L29). To make myself clear: I am not in general*

disapproving the conclusion of an important aqueous phase bbSOA fraction, but I would encourage a more critical assessment of the evidence provided, including a discussion of alternative processes that could at least in part explain the observations.

Authors Reply. In addition to the previous reply, partially answering also to this comment, authors would like to add few more specific considerations:

-HMSA is the product of sulfite complexation with formaldehyde, a reaction taking place only in aqueous-phase and inhibited by photochemistry. So the correlation of OOAx_BB-aq factors with this compound is in our opinion a clear evidence of the link with aqueous-phase chemistry and not only an indication of the preferential partitioning of formaldehyde at high ALWC conditions.

- we acknowledge that the AMS fragments and elemental ratios are not overwhelming evidences. For this reason we first showed that the fragments that were attributed to HMSA correlates with the HMSA concentration, quantified also with independent measurements (H-NMR). In addition, in our study the other fragments are used as ancillary evidences (for this reason reported in the Supplementary), not as the “main proof” that instead are the correlations with ALWC and HMSA and the variations with RH.

-we acknowledge also that the variability of AMS spectra is high. However literature indicates that this can be quite common looking secondary factors (OOAs) but in any case they can be used to investigate SOA sources. This is clearly visible in Tables S7 and S14 for example where the comparison between OOAs factor spectral profiles from SUPERSITO campaigns and many correspondent reference profiles from literature is reported.

-P5L18ff: *In its present form, the explanation of the alpha value approach is not really accessible to non-familiar readers. Is it important to be presented or would a reference to the method suffice?*

Authors Reply. Authors thank the Referee and re-phrase the explanation of the a-value approach, removing the equation and summarizing the concept as follows: “Similarly to the classical PMF solver (e.g., PMF2, PMF3, Paatero and Tapper, 1994), the ME-2 solver (Paatero, 1999) executes the positive matrix factorization algorithm. However, the user has the advantage to support the analysis by introducing *a priori* information, such as known factor profiles (FP), for example within the so-called *a*-value approach. The *a*-value is a scalar (defined between 0 and 1) that determines how much the resolved factor profiles are allowed to vary from the reference ones (Canonaco et al., 2013). For instance, applying an a-value of 0.05 lets $\pm 5\%$ variability to our FP solution with respect to the reference FP during the PMF iteration.”

-P5L15 and P5L22: *There is a contradiction in the range the alpha values can take.*

Authors Reply. Re-phrased and contradiction removed.

-P5L23: *“...represented by f.” What does f mean?*

Authors Reply. Sentence removed rephrasing the concept.

- P6L21: *“inorganic ions”?*

Authors Reply. Replaced as suggested

- P6L23: *ion charges are missing*

Authors Reply. Ion charges added as suggested

- P9L15: *“...in the summer period analyzed by itself” What does it mean? Please rephrase.*

Authors Reply. Rephrased removing “analyzed by itself”.

- P10L15: *Are local sources important for the urban increment, i.e. does wind direction play a role for the high variability?*

Authors Reply. In general, due to very low winds intensity (average values for all the campaigns of 1.89 ± 0.32 and 2.06 ± 0.17 m/s for BO and SPC respectively), the transport of pollutants along the two sites is supposed to be diffusive. So, the two sites are representative of two different background conditions (urban and rural), both poorly affected by really local sources in most of the cases (background sites) (http://www.arpae.it/sim/?osservazioni_e_dati/climatologia ; Ricciardelli et al., 2017).

The high variability affecting BBOA urban increment is something not expected, but (as expressed in the text, P10 L11, and showed in Section S2) probably it is due to the fact that the spring 2013 campaign was characterized by few nights with temperature colder (8°C) than the monthly average for May (18.5°C). These nights correspond to very sharp and intermittent peaks in BBOA concentrations (probably due to domestic heating active just for those few nights). For the rest, the BBOA concentrations during the campaign were really low (average values of 0.29 and $0.05 \mu\text{m}^{-3}$ at BO and SPC respectively), something not strange during May. So this very low and intermittent BBOA levels affect the urban increment value for this specific campaign and it is considered not representative.

- **P10L19:** *Is agricultural burning common in the area?*

Authors Reply. Actually agricultural burning is limited by the law: every Regional administration and sometimes also the different municipalities have a specific regulation on the topic especially during definite periods (e.g., dry summer, for fires prevention). However the open burning of agricultural waste is allowed for most of the year and continues to be a practice fairly common and difficult to regulate in rural areas (data confirmed by Regional Environmental Protection Agency of Emilia Romagna, ARPAE-ER). For this reason it is not possible to exclude the contribution of agricultural burning to the total BB aerosol burden at the rural background site (SPC), even if it is not possible to estimate its importance.

- **P12L21ff:** *I do not follow these conclusions. Both carboxylate as well as hydroxyl group formation can in principle take place both in the gas and in the aqueous phase. The references cited all relate to aqueous phase chemistry. Please elaborate. Also, in Fig. 6d it is the circles that show a negative slope and in contrast to what the text says in L21 and L23, these are the ones labelled "BB-aq".*

Authors Reply. Authors thank the Reviewer to highlight the discrepancies. We agree that in the present form the paragraph can be misleading for the reader. Our intention here is only to highlight the possible different oxidation pathways observed in the biomass burning ageing. Then in the subsequent paragraph we focus on the aqueous-phase chemistry. We rephrased the paragraph clarifying the message. We also add references of carboxylate formation not related to aqueous chemistry.

- **P13L6ff:** *Given the very different emission strengths of POA during a day, I doubt it can be used as a surrogate of the PBL. Also, which "above conclusion" (L9) is supported by similar variations and why?*

Authors Reply. Authors accept the Referee's suggestion: the sentence about POA as PBL surrogate is removed as well as the panel d) in Figures 7 and S5.

- **P15L14:** "ambient air"?

Authors Reply. Added

- **P15L24:** erase "extremely"

Authors Reply. Done

- **Fig. 6** is hard to digest. Please structure the legend according to the different campaigns and increase panel sizes of a) – d). Maybe make d) a separate figure to gain space for a) – c)? Check readability of greyish color (really “gold-green”?) Consider using a consistent color code for $OOAx_BB$ and $OOAx_BB-aq$. Right now, identifying the points in the scatter plot is a headache. Check the grey square for “no BB influence” for correct positioning. Text says 0.05 max, but the shaded area is < 0.05 in the plot.

Authors Reply. Done. A new version of the Fig. 6 is reported in the revised manuscript. The legend is now structured according to the different campaigns; the panel size is increased and the grey area for “no BB influence” is located in the right position. We prefer to leave the color distinction between different OOAs based on their O:C ratios and not on the formation process (which is instead highlight by the name of the factor). We prefer also to leave the panel d) in the same figure, even if increased in dimension.

- **Fig. 6a)** Some of the points are just at the edge of “BB influenced”, based on f60. Is this considered in data interpretation and what does it imply for their assignment as “ $OOAx_BB$ ”?

Authors Reply. As reported in the Supplementary section S2.2.3 (Validation of by Biomass Burning influenced OOAs) we performed additional tests in order to validate the attribution of the $C_2H_4O_2^+$ fragment (corresponding to the f60) to the OOA factors.

In the end, factors considered as $OOAx_BB$ are only those for which both average values and error bars (representing the standard deviation of all the additional tests performed) are located out of the gray shaded area indicating no influence of biomass burning (Fig. S3).

- **Fig. 8:** Check formatting of unit in axis labels.

Authors Reply. Done.

Replies to Referee #2

The authors would like to thank Anonymous Referee #2 for his/her comments.

The Referee's comments (in *italic*) followed by our replies are listed below.

-I must say here that the uploaded text font of the ACP manuscript is too small read offline. I suggest authors to take care of this part when uploading the revision.

Authors Reply. ACP has specific rules about the submission format (e.g., font size, line numeration, etc.). The Authors simply followed these rules to satisfy the requirements of the journal. The same is applicable for the last comment of the Referee #2: *Please provide line numbers continuously and also increase the font size of the text so that it will help us to properly evaluate the manuscript.*

-Grey shades in figures should be in 'black' with increase in font size for all the figures.

Authors Reply. We apply all the Referee's suggestions in the revised version of the paper.

-P2 L13: Is it 400 or 400,000 premature deaths??

Authors Reply. We thank the Referee to have notice the misspelled number. The correct number is four-hundred-thousand (400,000) of course. In order to avoid any confusion we completely remove the separator (putting 400000) in the revised version.

-P2 L20: replace 'proved' with 'established'

Authors Reply. Done.

-P3 L16: Aerosol Chemical Monitor (ACSM)??

Authors Reply. We thank the Referee to have notice the misspelling. The correct definition - Aerosol Chemical Speciation Monitor (ACSM) – is replaced in the revised version.

-P9L25-29: I could not follow the logic of arguments here? Do authors mean HOA are embedded/occluded in water-soluble OA components at SPC, which are scavenged by the fog and left behind the HOA, thus increasing the fossil-based emission contribution to SPC?? Some additional explanation is needed here.

Authors Reply. This behavior is better discussed in a previous publication by Gilardoni et al. (2014), specifically focused on SPC fall2011 campaign and cited at P9, L29. In the revised version we add a clearer linkage to this paper and we re-phrase the paragraph in order to clarify the concept, Here we just want to quickly explain the peculiarity of this specific campaign, representative of irregular meteorological conditions, and so not completely comparable with the others in term of relative contributions of the different OA components.

We would like to clarify that during the fall2011 campaign SPC was affected by persistent fogs during 14 days (out of the total 17). And (as reported in Gilardoni et al., 2014) fog scavenges aerosol components selectively, based on their solubility. For this reason, after fog formation, the interstitial aerosol (i.e., the fraction of the aerosol that is not scavenged and was measured in particle-phase) was dominated by particles enriched in carbonaceous aerosol, mainly black carbon and water-insoluble (or poorly soluble) organic aerosol. Moreover, analyzing the functional group composition and OA elemental ratios, Gilardoni et al. indicated that more oxidized OA was scavenged more efficiently than less oxidized OA.

HOA is the less oxidized (O:C ratio = 0.29) and less soluble fraction of OA, so it is likely the one that is scavenged less efficiently and therefore its fractional contribution in the interstitial aerosol increases. So, the aerosol composition and concentrations for this campaign at SPC referred to a

mixture of total OA and interstitial OA in fog conditions not fully comparable with the other campaigns.

-From Table 1, it is apparent that OA contribute almost 50% at both sampling sites (BO and SPC).

Authors Reply. Even if the comment is not clear, we think it refers to our sentence at P4, L22-23 “the organic aerosol (OA) component that represents the major fraction of submicron particles for most of the campaigns”. We don’t see any problem in this sentence: even when the OA fraction is less than 50% of the total mass there’s not any other single component (NO₃, SO₄, NH₄, etc.) accounting for a higher fraction of the PM₁ mass. To clarify, we modified the sentence as follows: “the organic aerosol (OA) component that represents the most abundant fraction of non refractory submicron particles mass for most of the campaigns...”

-It is bit confusing to see some places OOA and other places as SOA. Please maintain consistency throughout the manuscript.

Authors Reply. We partially accept the suggestion of the Reviewer. In the revised manuscript we use always “OOA” when we are talking about the PMF factors. However sometimes we keep also the “SOA” abbreviation when we refer to more general arguments, like potential sources and formation pathways.

-In Table 2, I understand the reason of HOA share decrease between BO and SPC. But the BBOA component show more at BO site compared SPC during spring 2013 but also somewhat higher or comparable for other seasons too. Some explanation is need in the manuscript.

Authors Reply. Unfortunately we don’t have enough reliable data to indicate a clear trend for BBOA concentrations between urban and rural site in all the seasons. The campaigns carried out in parallel at the two sites were only 4 (summer 2012, spring 2013 and fall 2011 and 2013). Among these: during the summer 2012 PMF didn’t identify any BBOA (no domestic heating during summer); the fall 2011 had its meteo peculiarity linked with fog events (already discussed in a previous reply here above and in the revised text at P9, L21-26) and so it is considered not comparable; the spring 2013 was characterized by very low and intermittent BBOA levels (as better discussed in some other replies below) and so it is not clear how it is representative of background conditions. The only campaign reliable to say something about differences in BBOA concentrations between urban and rural site is the fall 2013, suggesting “a higher contribution of BBOA in the rural areas, probably due to the more spread use of fire-places and wood-stoves for domestic heating and to additional possible sources, such as agricultural burning”, as clearly stated in the text.

-P10 L8: Is it because of the differences in the ambient temperature and photochemical activity between winter and summer controls their abundance whether it is NH₄NO₃ in winter/fall vs. (NH₄)₂SO₄ in summer and, hence, their correlation with OOA component. Add some additional explanations here.

Authors Reply. We believe that there re two possible explanations. As suggested by the Reviewer, the differences in temperature and relative humidity between winter and summer shift the partitioning of nitrate toward gas-phase (due to its volatility) during warm season. In addition, the different correlation suggests the possibility of a different oxidation pathways in secondary species formation between cold and warm season: a pathway characterized by cold temperature and high relative humidity (dominated by aqueous-phase processing and correlating more with nitrate) and another one more related with higher temperature and photochemical activity (correlating more with sulfate). We add in the revised manuscript a short explanation introducing the subsequent sections in which we developed more the concept.

-P10 L10: *Instead of calculating based on the overage, I recommend authors' to show the ratio of each fraction of OA between BO and SPC based on the box plots. This will give us a brief idea about the relative increment of emissions/formation processes contributing to observed compound classes of OA between both sites.*

Authors Reply. A more rigorous assessment of emission/formation processes would need a more detailed statistical treatment of the data variability between urban and rural sites, which is beyond the scope of the present manuscript. "Urban increment" here is not intended as a source apportionment tool but only as an instrument to discuss the differences observed between the urban and the rural sites. For this reason, even if we acknowledge the limitations of our urban increment assessment, we believe the calculation based on the average values is enough to have a rough idea of the most evident differences between the sites.

In any case, we calculated the same ratios using Median value (reported in the table here below), showing not substantial differences with Table 3. The high value corresponding to Spring 2013 (even higher than in Table 3) is consistent with the idea of few and intermittent high BBOA spikes, better discussed in the subsequent reply to the comment on Table 3.

Urban Increment using Median values	HOA	BBOA	OOA	OA TOT
SPRING 2013	1.73	8.85	1.28	1.42
SUMMER 2012	2.94		1.60	1.64
FALL 2013	1.70	0.64	1.20	1.11

-P10 L13: *Authors mentioned previously that in summer traffic is less at BO because of the shutdown of schools and public institutions, in which case, why the HOA fraction increased over BO compared to SPC in summer.*

Authors Reply. As the Urban increment clearly show (Table 3 and also previous comment), the HOA trend is constantly shifted toward higher concentrations at the urban site (BO) with respect to the rural one (SPC). This is quite expectable and already mentioned in the text: the urban site is closer to the HOA primary sources (traffic) with respect to the rural site and for this reason HOA concentrations are always higher at BO than at SPC, even when (during summer) are depleted (compared to the winter) due to the activities shutdown and increased dilution.

-In Table 3, why BBOA fraction is almost 6 fold higher at BO (urban) compared to the SPC (rural). This implies there exists a very strong local source of biomass combustion at BO compared to SPC, please clarify. Higher share of BBOA(%) over SPC in fall, why not is the case for winter or other seasons?

Authors Reply. Overall both the measurement sites are representative of background conditions: BO is representative of urban background, while SPC of the rural background. This because of two main reasons: 1- both the field stations are located quite far from any kind of direct local source and 2- due to very low winds intensity (average values for all the campaigns of 1.89 ± 0.32 and 2.06 ± 0.17 m/s for BO and SPC respectively), the transport of pollutants is supposed to be diffusive in the studied area (http://www.arpae.it/sim/?osservazioni_e_dati/climatologia ; Ricciardelli et al., 2017).

The high BBOA urban increment during spring 2013 is something not expected, but (as expressed in the text, P10 L11, and showed in Section S2) probably it is due to the fact that the spring 2013 campaign was characterized by few nights with colder temperature (8°C) then the monthly average for May (18.5°C). These nights correspond to very sharp and intermittent peaks in BBOA concentrations (probably due to domestic heating active just for those few nights). For the rest, the BBOA concentrations during the campaign were really low especially in SPC (average values of

0.05 μm^{-3}), something not strange during May. So this very low and intermittent BBOA levels affect the concentrations for this specific campaign, which is considered for this reason not representative.

About winter seasons: unfortunately our dataset is not comprising a winter campaign carried out in parallel at both the sites and for this reason we can not argue nothing about the urban increment during winter.

-P11 L6: Why focus only on these two factors?

Authors Reply. Considering the big amount of data already reported in the paper, the Authors decided to focus more only on the two most innovative aspects emerged from this study, namely the biomass burning influence on SOA components and the aqueous-phase processing affecting some of these components. These two aspects are represented by the OOA_BB factors deeper examined, indeed.

-P11 L30-31: How this fraction of OOA_BB was estimated here?

Authors Reply. This fraction is represented by the mass contribution of the OOA factors which, looking the f60 vs f44 space, resulted influenced by biomass burning.

As reported in the Supplementary section S2.2.3 (Validation of by Biomass Burning influenced OOAs) we performed additional tests in order to validate the attribution of the $\text{C}_2\text{H}_4\text{O}_2^+$ fragment (corresponding to the f60) to the OOA factors.

In the end, factors considered as OOA_x_BB are only those for which both average values and error bars (representing the standard deviation of all the additional tests performed) are located out of the gray shaded area indicating no influence of biomass burning (Fig. S3).

-Figure 6 panel resolution and font sizes need to be improved? This figure is not readable at all offline. Grey shade text in panel d should be converted to black.

Authors Reply. A new version of the Fig. 6 is reported in the revised manuscript.

-P12 L19-20: These sentences are not clear, please rewrite. The slope line between triangles and circles seems to be zero (i.e., OOA_x_BB-aq) and those between triangles and squares (OOA_x_BB) is like between -0.5 and one. P12 L24: This is contradicting the above classification on L19-20. Please check.

Authors Reply. Authors thank the Reviewer to highlight the discrepancies. The text was mixed up during writing and so the description of the figure is not consistent. We rephrased the paragraph clarifying the message.

-P12 L29: What are the input parameters to ISORROPIA-II, which mode is used, please provide.

Authors Reply. Info provided in section 2.3 of the revised manuscript.

-P13 L7: sentence should read like this, 'Dividing the individual OOA fractions with the total POA'

Authors Reply. Authors, accepting a comment from Referee #1, removed the sentence about POA as PBL surrogate and removed also the panel d) in Figures 7 and S5.

-In Figure 8, what is the OOA2, OOA3, OOA refers to, Please clarify.

Authors Reply. OOA factors are numerically ordered based on their O:C ratios, independently on the influence of a specific source, as explained in the text (P11,L5) and in the captions of Figure 6, showing for the first time in the paper the different OOAs. To improve clarity we add this information in the revised manuscript also in the caption of Figure 8.

-P13 L16: I can see m/z 29 signal but not 58 from figure 8 (left panel). Did I miss something here?

Authors Reply. The ion at m/z 58.01 01 ($C_2H_2O_2^+$) is associated in literature to aqueous-phase reactions because is one of the typical fragments of precursors of SOA via cloud processing, like methylglyoxal and glyoxal (Carlton et al., 2007; Altieri et al., 2008). We state this clearly in the text (P13, L27-29 of revised manuscript).

Since m/z 58.01 has much lower fractional abundances (f58) with respect to other fragments in the spectra this ion is not easily distinguishable in Figure 8 (where the spectra are reported in a very synthetic way). We acknowledge this but we prefer to avoid adding more information in Figure 8 (already quite packed).

Anyway the f58 is always higher in all the OOA_BB-aq with respect to the others OOA_BB factors (as showed in the additional graph here below). This feature, associated with the abundance of m/z 29 (CHO^+), can help in the identification of the BB-aqSOA factors, even if it is not a sufficient proof (because the correlation with ALWC and HMSA together with the study of the variations with RH are more important to unambiguously identify those factors).

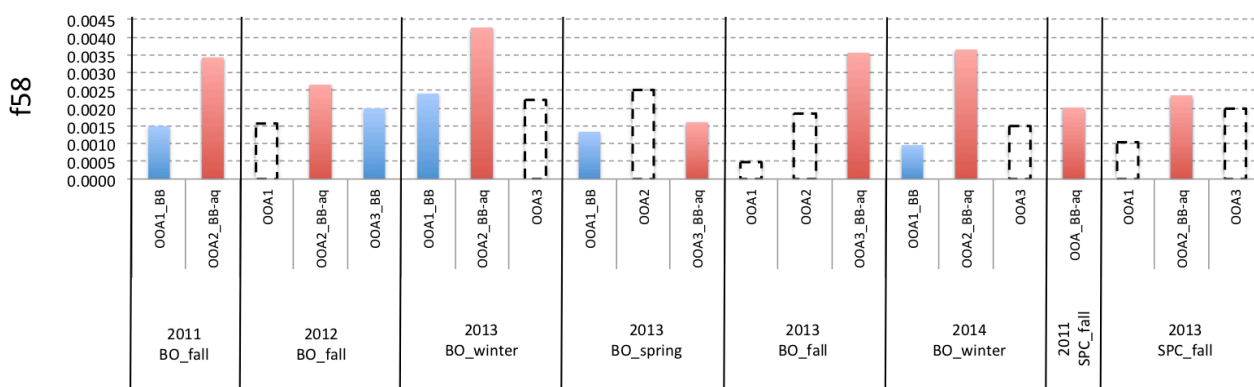


Figure AR1. Fractional abundance of m/z 58.01 (f58) in the spectral profiles of each OOA factor identified by PMF analysis in all the campaigns where a BB-aqSOA factor is identified. Blue bars represent the OOA_BB not associated with aqueous-processing; Red bars are the OOA_BB-aq factors; the bars with dashed line represent all the other OOAs (not influenced by biomass burning).

-P13 L21: mention those specific fragment ions here within parenthesis.

Authors Reply. They are mentioned in the subsequent sentences. The Authors don't see the need to add them also before within parenthesis.

-In Figure 9, why there is no such presence of aq-SOA despite more sunlight and having precursors at both sites. Some explanations needed in the manuscript.

Authors Reply. The question is hard to understand because no specific reference to the text or the figure is provided. Anyway, we think the comment is referring to the absence of a BB-aqSOA at SPC during Spring 2013. We believe it is probably related to the very low and intermittent BBOA levels affecting this specific campaign (already discussed above) and not allowing an important BB-SOA formation.

List of relevant changes

Relevant changes made in the initial Manuscript (as already discussed in the Replies to the Referees and tracked in the following marked-up version) are listed below.

- **Section 2.2, P5, L13-18** (of Revised version): Following the suggestion of Referee #1, the Authors re-phrase the explanation of the *a*-value approach, removing the equation and summarizing the concept as follows: “Similarly to the classical PMF solver (e.g., PMF2, PMF3, Paatero and Tapper, 1994), the ME-2 solver (Paatero, 1999) executes the positive matrix factorization algorithm. However, the user has the advantage to support the analysis by introducing *a priori* information, such as known factor profiles (FP), for example within the so-called *a*-value approach. The *a*-value is a scalar (defined between 0 and 1) that determines how much the resolved factor profiles are allowed to vary from the reference ones (Canonaco et al., 2013). For instance, applying an *a*-value of 0.05 lets $\pm 5\%$ variability to our FP solution with respect to the reference FP during the PMF iteration.”

- **Section 3.2, P9, L23ff** (of Revised version): following the suggestions of Referee #2, the Authors added some more explanation about the SPC 2011 fall campaign.

- **Section 3.2, P10, L3-9** (of Revised version): as requested by Referee #2, the Authors clarify their idea about the possible reasons of a different correlation between ammonium nitrate and/or sulfate with OA. We added to the text this sentence: “This behavior likely reflects the differences in temperature and relative humidity between winter and summer, which shift the partitioning of nitrate toward gas-phase (due to its volatility) during warm season. In addition, the different correlation 5 suggests the possibility of a different oxidation pathways in secondary species formation between cold and warm seasons: a pathway characterized by cold temperature and high relative humidity (correlating with nitrate) and another one related with higher temperature and photochemical activity (correlating more with sulfate). The latter hypothesis will be better developed in the following sections.”

- **Section 4.2, P12, L20-28** (of Revised version): following the suggestions of Referee #1 and #2, the Authors rephrased the paragraph removing the highlighted discrepancies and clarifying the message of possible different oxidation pathways observed in the biomass burning ageing.

- **Section 4.3, P13, L4-11** (of Revised version): in order to address the main concerns of Referee #1, we added a sentence to clarify the message: “We suggest here that formaldehyde (as well as a number of other gaseous compounds including ketones, aldehydes, and small carboxylic acids in the BB plumes, Schauer et al, 2001; Andreae, 2019) would preferentially partition into particles at high ALWC and would react to form HMSA (and/or other products). Then the products of these aqueous phase reactions (such as HMSA) remain in the particle phase after water evaporation, changing the chemical composition of the organic aerosol. For this reason, considering also that HMSA formation is inhibited by photochemistry (due to its fast reaction with ozone) and that the analyzed aerosol was dried before sampling, the correlation of some factors with HMSA can be considered a reliable evidence of aqueous-phase formation pathway of some OA fractions.”

- **Figures 6, 7 and 8** were modified according with the suggestions of Referee #1 and #2.

The impact of biomass burning and aqueous-phase processing on air quality: a multi-year source apportionment study in the Po Valley, Italy

M. Paglione¹, S. Gilardoni¹, M. Rinaldi¹, S. Decesari¹, N. Zanca^{1,*}, S. Sandrini¹, L. Giulianelli¹, D. Bacco², S. Ferrari², V. Poluzzi², F. Scotto², A. Trentini², L. Poulain³, H. Herrmann³, A. Wiedensohler³, F. Canonaco⁴, A. S. H. Prévôt⁴, P. Massoli^{5,6}, C. Carbone⁷, M.C. Facchini¹, S. Fuzzi¹

¹Italian National Research Council - Institute of Atmospheric Sciences and Climate (CNR-ISAC), Bologna, 40129 Italy

²Regional Agency for prevention, environment and energy (ARPAE) of Emilia-Romagna, Bologna, Italy

³Leibniz-Institut für Troposphärenforschung (TROPOS), Leipzig, 04318, Germany

⁴Laboratory of Atmospheric Chemistry, Paul Scherrer Institute, Villigen PSI 5232, Switzerland

⁵Aerodyne Research, Inc. Billerica, MA, USA

⁶MultiSensor Scientific, Inc., Greentown Labs, Sommerville, MA, USA

⁷Proambiente S.c.r.l., CNR Research Area, Bologna, Italy

*now at Department of Chemistry and Institute for Atmospheric and Earth System Research (INAR), University of Helsinki,

FI-00014, Finland

Correspondence to: Marco Paglione (m.paglione@isac.cnr.it)

Abstract.

The Po Valley (Italy) is a well-known air quality hotspot characterized by Particulate Matter (PM) levels well above the limit set by the European Air Quality Directive and by the World Health Organization, especially during the colder season. In the framework of the Emilia-Romagna regional project "SUPERSITO", the southern Po Valley submicron aerosol chemical composition was characterized by means of High-Resolution Aerosol Mass Spectroscopy (HR-AMS) with the specific aim of organic aerosol (OA) characterization and source apportionment. Eight intensive observation periods (IOPs) were carried out over four years (from 2011 to 2014) at two different sites (Bologna, BO, urban background and San Pietro Capofiume, SPC, rural background), to characterize the spatial variability and seasonality of the OA sources, with a special focus on the cold season.

On the multi-year basis of the study, the AMS observations show that OA accounts for an average $45\pm 8\%$ (ranging 33-58%) and $46\pm 7\%$ (ranging 36-50%) of the total non-refractory submicron particle mass (PM₁-NR) at the urban and at the rural site, respectively. Primary organic aerosol (POA) comprises biomass burning ($23\pm 13\%$ of OA) and fossil fuel ($12\pm 7\%$) contributions with a marked seasonality in concentration. As expected, the biomass burning contribution to POA is more significant at the rural site (urban/rural concentrations ratio of 0.67), but it is also an important source of POA at the urban site during the cold season, with contributions ranging from 14 to 38% of the total OA mass.

Secondary organic aerosol (SOA) contribute to OA mass to a much larger extent than POA at both sites throughout the year ($69\pm 16\%$ and $83\pm 16\%$ at urban and rural, respectively), with important implications for public health. Within the secondary fraction of OA, the measurements highlight the importance of biomass burning ageing products during the cold season, even

at the urban background site. This biomass burning SOA fraction represents 14-44% of the total OA mass in the cold season, indicating that in this region a major contribution of combustion sources to PM mass is mediated by environmental conditions and atmospheric reactivity.

Among the environmental factors controlling the formation of SOA in the Po Valley, the availability of liquid water in the aerosol was shown to play a key role in the cold season. We estimate that organic fraction originating from aqueous reactions of biomass burning products ("bb-aqSOA") represents 21% (14-28%) and 25% (14-35%) of the total OA mass and 44% (32-56%) and 61% (21-100%) of the SOA mass at the urban and rural sites, respectively.

1 Introduction

10 Ambient air pollution represents the highest environmental risks for human health, leading to about 3 million premature deaths every year (WHO, 2016) due to the exacerbation of respiratory and cardio-vascular diseases especially in young kids and elderly people. In Europe atmospheric pollution is responsible for more than 400000 premature deaths a year (EEA, 2016), with the largest share due to fine particulate matter (PM_{2.5} and PM₁) exposure. Organic aerosol (OA) accounts for 20 to 90% of fine particle mass worldwide (Zhang et al, 2007), and for up to 50% (20-90%) of fine particle mass in Europe 15 (Putaud et al., 2010). OA global budget and atmospheric processing are still characterized by large uncertainties (Hallquist et al., 2009). A better knowledge of OA is essential to support effective air quality control and remediation measures.

OA is directly emitted by various sources, including traffic, other combustion sources and biogenic emissions, and can also be produced via secondary formation pathways in the atmosphere (Hallquist et al., 2009). In particular, our understanding of the formation mechanisms and evolution processes of secondary OA (SOA) is still largely uncertain.

20 Direct quantification of SOA in the ambient aerosol is challenging, but many recent studies have established that oxygenated OA (OOA) determined by multivariate statistical analysis (e.g., positive matrix factorization, PMF) of OA fragmentation mass spectra is a good proxy of SOA (Zhang et al., 2007; Ulbrich et al., 2009). Therefore, OOA is widely used to study the abundance and formation mechanisms of SOA. Although several types of these OOAs were isolated in ambient aerosol everywhere (often representing more than half of the total OA (Zhang et al., 2007; Ng et al., 2010; Crippa et al., 2014)), their 25 link to a specific source or mechanism remains largely undetermined. This is a consequence of their complexity in terms of chemical and physical properties, and the difficulty of reproducing the real conditions in which SOAs are formed/transformed. As a result, traditional models often show substantial discrepancies in simulating SOA mass concentrations (Kleinman et al., 2008; Matsui et al., 2009) and oxidation states (Chen et al., 2011), especially in wintertime.

30 The Po valley, located in northern Italy, is amongst the most polluted areas in Europe (EEA, 2016). It is surrounded by the Alps to the North and North-West and by the Apennines to the South. The occurrence of frequent and prolonged low-wind periods and atmospheric stability conditions favour the accumulation of particulate and gaseous pollutants locally emitted,

Marco Paglione 31/7/y 14:12

Eliminato: .

Marco Paglione 31/7/y 14:31

Eliminato: proved

especially during the cold months. The distinctive features of the Po Valley make it an interesting “laboratory” to study the development of POA and SOA concentrations in the ambient atmosphere.

The SUPERSITO project (www.arpae.it/supersito) is a comprehensive study of atmospheric particulate matter pollution in the Emilia-Romagna Region, encompassing the southern part of the Po Valley from the Po river to the Apennines. Overall, the project deals with chemical, physical and toxicological parameters of the aerosol and integrates them with epidemiological and medical assessments through interpretative models. Results about aerosol chemical characterization using offline techniques were presented by Ricciardelli et al. (2017).

Here we describe the results of HR-AMS PM₁ measurements carried out during eight intensive measurement campaigns with a focus on OA source apportionment. Previous projects have investigated the properties of fine aerosols at urban, rural and regional sites of the Po valley, including their chemical features (Carbone et al., 2014; Putaud et al., 2002, 2010; Saarikoski et al., 2012), and main sources (Belis et al., 2013; Gilardoni et al., 2011; Larsen et al., 2012; Perrone et al., 2012). Further studies based on aerosol mass spectrometer (AMS) measurements have been conducted in the same area during specific field experiments with the aim of characterizing specific phenomena and seasonal features (e.g. fog events, cooking aerosols, biomass burning emissions, etc.) (Gilardoni et al. 2014; Decesari et al., 2014; Paglione et al., 2014; Dall’Osto et al., 2015). Nevertheless, systematic AMS observations in the Valley are available from very few studies. Bressi et al. (2016) using a 1-year long dataset of measurements by an Aerosol Chemical Speciation Monitor (ACSM) described the chemical composition and the organic PM₁ sources of the north-west edge of the Po Valley at the rural background site of Ispra, 60 km northwest of Milan.

In this study, we analyze a multi-year dataset of high resolution measurements carried out at two different sites (Bologna and San Pietro Capofiume) exploring for the first time the spatial-temporal variability of OA sources, chemical features and formation/transformation processes in the southern part of the Po Valley. A special focus is dedicated to the interpretation of the main sources and formation/transformation processes of the SOA in the region active during the cold period.

2 Material and methods

2.1 Measurement field campaigns

Eight intensive observation periods (IOPs) were carried out over four years (from November 2011 to June 2014) at two different sites of the southern part of the Po Valley (Bologna, BO, urban background and San Pietro Capofiume, SPC, rural background). Figure 1 reports a map of the measurement sites and a time-line of the field campaigns carried out during SUPERSITO project. Bologna is located at the foot of the Apennines and is an important population basin for the region (400,000 inhabitants), impacted by significant industrial and agricultural activities, and crossed by several major highways. The BO measurements site is located at the National Research Council (CNR) Research Area (N 44°31’29”, E 11°20’27”). The rural background station of San Pietro Capofiume (SPC) is located in a sparsely populated flat countryside (N

44°39'15", E 11°37'29") surrounded by kilometers of flat lands in the southeast part of Po Valley, 30 km north-east of Bologna and is representative of the regional background. This site is used for many atmospheric characterization studies and research projects (Saarikoski et al., 2012; Paglione et al., 2014; Decesari et al., 2014; Sandrini et al., 2016).

5 During the four-year project, the intensive campaigns were programmed to account for the marked seasonality in both sources and weather conditions of this region. Nevertheless, most of the SUPERSITO campaigns took place in the cold season (3 campaigns in fall and 2 in winter, out of 8 in total) when the highest PM levels are found. Similar to other continental sites, during fall-winter the reduced height of the Planetary Boundary Layer (PBL) and calm wind conditions favor the accumulation of pollutants and are responsible for the rise of PM concentration (Perrone et al., 2012; Stanier et al., 2012; Bressi et al., 2013). Another feature of the cold months in this area is the high relative humidity, which leads to fogs and hazes (i.e., conditions of high aerosol liquid water content, ALWC). The consequence of these meteorological conditions on PM concentrations is twofold: it promotes both wet removal (Gilardoni et al., 2014; Giulianelli et al., 2014; Montero-Martinez et al., 2014) and aqueous-phase processing with SOA formation (Gilardoni et al., 2016).

2.2 Aerosol mass spectrometer measurements and apportionment of organic fraction

15 During all of the SUPERSITO campaigns, the mass loading and the size-resolved chemical composition of submicron aerosol particles were obtained online by the Aerodyne high-resolution time-of-flight aerosol mass spectrometer (HR-TOF-AMS, Aerodyne Research, Canagaratna et al. 2007). The HR-TOF-AMS provides measurements of the non-refractory sulfate, nitrate, ammonium, chloride, and organic mass of the submicron particles (NR-PM1). The average concentrations of NR-PM1 chemical components and their relative contributions as measured by AMS in each campaign are reported in the Supplemental material (Table S1 and Figure S1). For some of the SUPERSITO campaigns, specific studies have already been published. We refer to Gilardoni et al. (2014 and 2016) for the SPC fall 2011 and BO winter 2013 campaigns, respectively, and to Sullivan et al. (2016) for the SPC summer 2012 campaign. In this paper we focus on the organic aerosol (OA) component that represents the most abundant fraction of submicron particles mass for most of the campaigns, ranging between 33 and 58% of NR-PM1 (concentration range: 1.8-18.4 $\mu\text{g m}^{-3}$), consistent with the value found by Jimenez et al. (2009), Ng et al. (2010) and Crippa et al. (2014). Table 1 summarizes the average OA concentration for each site and season and the relative organic contribution to the NR-PM1 as measured by the HR-TOF-AMS.

The working principle of the HR-TOF-AMS is described in detail in Canagaratna et al. (2007), Jayne et al. (2000), and Jimenez et al. (2003). Briefly, during all the campaigns, the HR-TOF-AMS was operating by alternating between "V" and "W" ion path modes every 5 min. The concentrations reported here correspond to the data collected in V mode. The resolving power (DeCarlo et al., 2006) of the V-ion mode was about 2000-2200 during all the campaigns.

Ionization efficiency (IE) calibrations were performed before and after every campaign, and approximately once a week during the campaigns. Filter blank acquisitions during the campaign were performed once a day to evaluate the background

Marco Paglione 5/8/y 15:35

Eliminato: major

and correct for the gas-phase contribution. All data were analyzed using the standard ToF-AMS analysis software SQUIRREL v1.51 and PIKA v1.10 (D. Sueper, available at: <http://cires.colorado.edu/jimenez-group/ToFAMSResources/ToFSoftware/index.html>) within Igor Pro 6.2.1 (WaveMetrics, Lake Oswego, OR). The HR-TOF-AMS collection efficiency (CE) was calculated based on aerosol composition, according to Middlebrook et al. (2012) and evaluated against parallel offline measurements (see section 2.3 and Table S2 in the Supplement). The aerosol was dried to about 35-40% by means of a Nafion drier before sampling with the HR-TOF-AMS.

The organic fraction (OA) measured by HR-TOF-AMS was apportioned using the Positive Matrix Factorization approach (PMF; Paatero and Tapper, 1994; Lanz et al., 2007; Ulbrich et al., 2009; Zhang et al., 2011) by applying the Multilinear Engine 2 solver (ME-2, Paatero, 2000) controlled within the Source Finder software (SoFi v4.8, Canonaco et al. 2013; Crippa et al., 2014).

Similarly to the classical PMF solver (e.g., PMF2, PMF3, Paatero and Tapper, 1994), the ME-2 solver (Paatero, 1999) executes the positive matrix factorization algorithm. However, the user has the advantage to support the analysis by introducing *a priori* information, such as known factor profiles (FP), for example within the so-called *a*-value approach. The *a*-value is a scalar (defined between 0 and 1) that determines how much the resolved factor profiles are allowed to vary from the reference ones (Canonaco et al., 2013). For instance, applying an *a*-value of 0.05 lets $\pm 5\%$ variability to our FP solution with respect to the reference FP during the PMF iteration.

The standardized source apportionment strategy introduced in Crippa et al. (2014) is systematically applied to the 12 available HR-TOF-AMS datasets (8 from BO and 4 from SPC), consisting of the organic mass spectra over time and the corresponding errors.

The interpretation of the retrieved source apportionment factors as organic aerosol sources is based on the comparison of their mass spectral profiles with reference ones (Table S5, S6 and S7), on the correlations with external data (see Table S8) and on the investigation of their diurnal trends. Details of the factor analysis (number of factors chosen, Q and residuals diagnostic plots, constrained factor profiles and *a*-values if applied) are reported for each campaign in the supplemental section S2.

2.3 Additional measurements and analytical techniques

Additional measurements from the routine daily program of the SUPERSITO project are used in this study as ancillary data. PM2.5 daily samples were collected by a low volume sampler (Skypost PM, TCR TECORA Instruments operated at the standard flow-rate of 38.3 L min⁻¹) on quartz fiber filters (PALL Tissu Quartz 2500 QAO-UP 2500 filters, 47mm) during all the project periods for the analysis of the carbonaceous fractions (total carbon, TC; organic carbon, OC; and elemental carbon, EC) by thermo-optical transmittance (Sunset, Laboratory Inc., Oregon, USA, using the EUSAAR2 thermal protocol,

Marco Paglione 25/7/ly 17:59
Eliminato: factor time series (FT) and/or

Marco Paglione 25/7/ly 18:09
Eliminato: (ranging

Marco Paglione 25/7/ly 18:10
Eliminato: from

Marco Paglione 25/7/ly 18:10
Eliminato: to

Marco Paglione 25/7/ly 18:04
Eliminato: values larger than

Marco Paglione 25/7/ly 18:11
Eliminato:)

Marco Paglione 25/7/ly 18:08
Eliminato: (($f_{i,c}$)solution, and ($g_{i,c}$)solution)

Marco Paglione 25/7/ly 18:08
Eliminato: input

Marco Paglione 25/7/ly 18:09
Eliminato: (f_i , g_i), as defined in Eq. (1)

Marco Paglione 25/7/ly 18:10
Eliminato: - ... [1]

Marco Paglione 25/7/ly 17:57
Eliminato:)

Marco Paglione 25/7/ly 18:11
Eliminato:)

Marco Paglione 25/7/ly 18:12
Eliminato: In our work we only constrained the mass spectra represented by f_i -

Marco Paglione 25/7/ly 17:56
Formattato: Tipo di carattere:Corsivo

Cavalli et al., 2010; Ricciardelli et al., 2017) and of polar organic compounds (anhydrosugars and acids) by GC/MS analysis (Pietrogrande et al., 2014). Due to the elevated PM loading during the first experiment in fall 2011, the discrimination between OC and EC was not possible for the filters collected and only TC data are available for that specific campaign.

5 Black carbon (BC) was calculated from aerosol absorption coefficient measurements (when available) by a single-wavelength (573 nm) and a multi wavelength (467, 530, and 660 nm) Particle Soot Absorption Photometer PSAP (Bond et al., 1999), as previously described (Gilarioni et al., 2011; Gilarioni et al., 2016; Costabile et al., 2017).

Size-segregated aerosol particles were also sampled by a Berner impactor (flow rate 80 L min⁻¹) (Matta et al., 2003). The
10 Berner impactor collects particles on five stages, corresponding to the following particle aerodynamic diameter cutoffs: 0.14, 0.42, 1.2, 3.5, and 10 μm. Sampling was performed continuously during the intensive campaigns. Each day we collected two samples: a daytime sample (from ≈ 09:00 to 17:00 LT during fall/winter, and from ≈ 9:00 to 21:00 LT during spring/summer), and a night-time one (from 17:00 to 09:00 LT during fall/winter, and from 21:00 to 09:00 during spring/summer). Particles collected were extracted in water and analyzed by means of evolved gas analysis and ion
15 chromatography for quantification of the water-soluble Total carbon (TC) and the inorganic ions. Elemental and chromatographic analyses of the filter samples are used to validate the AMS data for the main aerosol components (Org, NO₃⁻, SO₄²⁻, NH₄⁺ and Cl⁻) and PMF factors, as reported in the Supplemental (Table S2 and Table S8).

Submicron particles were also sampled on prewashed and prebaked quartz-fiber filters (PALL, 9 cm size) using HiVol
20 samplers (a dichotomous sampler Universal Air Sampler, model 310, MSP Corporation at a constant nominal flow of 300 L min⁻¹ or, alternatively, a TECORA eco-highvol equipped with Digitel PM1 sampling inlet, nominal flow 500 L min⁻¹) located at ground level. Typically, two filters were collected every day in parallel with the Berner impactor sampling time. The HiVol quartz-fiber samples were analyzed to identify organic molecular tracers (e.g., levoglucosan, hydroxymethanesulfonate (HMSA) and low-molecular weight amines) using proton nuclear magnetic resonance (¹H-NMR)
25 spectroscopy according to Decesari et al. (2006). The concentration of the organic tracers identified by NMR are correlated with the PMF-factors identified by the AMS, trying to detail their chemical features and infer their sources and atmospheric processing (especially for the OOA).

Meteorological data are provided by the Hydro-Meteo-Climate Service of the Regional Environmental Protection Agency of
30 Emilia Romagna (ARPAE). In addition, aerosol liquid water content that is associated with the aerosol inorganic species (K⁺, Ca²⁺, Mg²⁺, NH₄⁺, Na⁺, SO₄²⁻, NO₃⁻, Cl⁻) was predicted by the ISORROPIA-II model [used in reverse mode](#) (Fountoukis and Nenes, 2007).

Marco Paglione 25/7/y 18:21

Eliminato: d

Marco Paglione 25/7/y 18:23

Eliminato: species

Marco Paglione 25/7/y 18:20

Formattato: Apice

Marco Paglione 25/7/y 18:20

Formattato: Apice

3 Results and discussion

3.1 Organic aerosol source apportionment

The source apportionment procedure allowed the identification of various components tracing the contributions of primary and secondary organic aerosol sources: hydrocarbon-like organic aerosol (HOA) resulting from the combustion of fossil fuels (e.g., vehicular traffic); BBOA (biomass burning organic aerosol) resulting from biomass combustion, mainly associated to wood combustion for domestic heating; COA (cooking organic aerosol) associated with specific food cooking practices. The latter is found just as a minor component of OM and only in one campaign at BO (spring 2014). The rest of the mass of sub-micrometer organic aerosol consists of oxygenated organic aerosols (OOAs), representative of secondary formation and/or ageing processes in the atmosphere. Factor analysis extracted different types of OOA with distinct time trends and/or spectral features. In this section, we will consider the OOA factors as a whole, while in Section 4 we will discuss a source attribution for the individual factors.

Figure 2 shows the average mass spectra of all the identified HOA (n=12), BBOA (n=10) and COA (n=12) (reduced from high resolution, HR, to unit mass resolution, UMR, for better readability) together with their standard deviation. The comparison between our profiles from the Po Valley and reference profiles is reported in supplemental section 2.1 in term of theta-angle (θ) between the spectra (Kostenidou et al., 2009). The theta-angle is a metric for the similarity between two spectra ($\theta < 15^\circ$ good; $15^\circ < \theta < 30^\circ$ partial; $\theta > 30^\circ$ bad similarity).

The HOA profile is characterized by peaks corresponding to aliphatic hydrocarbons including m/z 27, 41, 43, 55, 57, 69, 71, etc. (Canagaratna et al., 2004). The median HOA profile in our study shows a good overlap (mostly $\theta < 15^\circ$) with almost all the reference spectra compared, as expected for this type of source which is quite reproducible in terms of AMS spectral characteristics (Crippa et al., 2014). Among the HOA profiles found for the individual campaigns, only one (SPC fall 2011) shows low correlations with the others from this study and with the references. Such discrepancy must be due to the peculiar conditions during the campaign, as the numerous fog events strongly impacted the OA time trends and, in turn, also the ability of PMF to resolve sources profiles. The aerosol observations during the SPC fall 2011 campaign have been already thoroughly described by Gilardoni et al. (2014) and will be summarized later in the discussion.

Unlike the HOA, the BBOA profiles are more variable, in agreement with earlier findings (Grieshop et al., 2009; Heringa et al., 2011) showing that the biomass burning aerosol mass spectrum is strongly affected by burning conditions and types of wood/biomass. Nonetheless, the deconvolved BBOA profiles show good similarities with many reference spectra from previous studies with their characteristic peaks at m/z 29 (CHO^+), 60 ($\text{C}_2\text{H}_4\text{O}_2^+$) and 73 ($\text{C}_3\text{H}_5\text{O}_2^+$), which are associated with fragmentation of anhydrosugars such as levoglucosan (Alfarra et al., 2007; Aiken et al., 2009).

The COA factor was identified without any constrain only during the BO spring 2014 campaign. Its spectral profile exhibits good similarities with the correspondent reference spectra (Mohr et al., 2012; Crippa et al., 2013a). The presence of this

Marco Paglione 31/7/y 19:01
Eliminato: S

COA factor reduced sensibly the model residuals in the central part of the day and it is therefore considered in the final solution.

The more oxidized factors (OOA) differ from each other for the fractional abundance of m/z 43 and 44 and for the intensity of other fragments such as 29, 60 and 73. The spectral characteristics of the specific OOA factors are discussed in Section 4.

5

The correlation parameters between the time trends of AMS organic factors and of atmospheric tracer compounds are reported in Table S8. The time series of HOA correlates with that of elemental carbon (EC) or black carbon (BC) and with that of NO_x. The correlation with NO_x points to major sources of HOA from traffic. The trend of BBOA concentrations instead correlates with the trend for levoglucosan (measured by off-line techniques: GC/MS or ¹H-NMR) and with the organic fragments at m/z 60 and 73, which have been previously shown as good markers for biomass burning (Alfarra et al., 2007; De Carlo et al., 2008; Aiken et al., 2009). The concentration ratios between POA factors and tracer compounds (e.g., HOA/BC, HOA/NO_x, BBOA/levoglucosan, etc.) are reported in Table S9 and compared with literature ranges. The overall good agreement between these source-specific ratios and the literature ranges confirms our apportionment of POA components. The time trends of the OOA concentrations are contrasted with those of secondary inorganic species (i.e. NO₃⁻, SO₄²⁻ and NH₄⁺) and with the organic fragments at m/z 43 (Org_43 = C₂H₃O⁺) and 44 (Org_44 = CO₂⁺) generally exhibiting good correlations.

10

15

The identified factors daily trends (HOA, BBOA, COA and OOA) are shown in Figure 3. Median diurnal patterns are reported together with the 10th, 25th, 75th and 90th percentiles for each factor, for the lumped datasets from all SUPERSITO campaigns and separately for Bologna (BO) and San Pietro Capofiume (SPC).

20

The daily trends of each organic component exhibit consistent characteristics during all the campaigns. HOA presents a diurnal cycle characterized by two maxima corresponding to the rush hours (impacted by the greatest vehicular traffic) between 8-9 and 18-20, in agreement with the attribution of this fraction to traffic sources. This is especially evident at the urban site of Bologna compared to the rural one in which the concentrations of HOA are lower and rush hour signatures are weak, as expected for a rural background site. BBOA is characterized by a daily cycle with a midday minimum and a nighttime maximum. This behavior reflects the combination of two factors: the influence of the mixing layer height - which favors pollutant accumulation near the ground at nighttime - and the daily pattern of the emissions from domestic heating, increasing in the evening/night hours. The concentrations of COA exhibits a characteristic daily trend with two maxima corresponding to the hours of main meals, one in the central hours of the day (12-14) and the other in the evening (20-21, more pronounced due to the shallow boundary layer after the sunset). Finally, OOA exhibits an almost flat daily trend, reflecting its regional nature or the influence of multiple secondary formation processes. Therefore, the weak diurnal trends of OOA were not informative of potential sources of SOA in this region.

25

30

Marco Paglione 31/7/y 19:01

Eliminato: S

3.2 POA and SOA contributions, seasonality and spatial variability

Table 2 summarizes the site-specific and campaign-specific contributions of OA components determined by AMS factor analysis (see also Figure 4). A few clear seasonal patterns can be identified especially for the Bologna urban site for which a higher number of measurements are available (Figure 4).

5 In Bologna, HOA contributes for 11-18% of the mass of sub-micrometric OA in fall-winter and for 6-12% in spring-summer. The slightly lower average HOA contribution during warmer season likely reflects the combination of two aspects: the reduction of work and school activities in summertime nearby the sampling areas, leading to a reduction of traffic emissions, and a possible meteorological effect due to the higher mixing-layer, resulting in an enhanced dilution of the primary pollutants locally emitted.

10 The contribution of BBOA varies instead from 17-38% in the fall-winter campaigns to 0-14% in summer-spring. In particular, the contribution of BBOA has not been detected in the summer period. Biomass burning therefore dominates over fossil fuel combustion as a source of primary organic aerosols at the urban site during the cold season. At the same site, OOA contributes for 44-68% of the mass of sub-micrometric OA in fall and winter, while its contribution in spring and summer period increases to 74-92%. The higher relative contribution of OOA in the warm period is expected given the
15 reduction of residential combustion and the increased photochemistry. However, the OOA fraction in the cold season is still quite high, considering the latitude and climate of Bologna, where sunshine duration in winter is less than 3 h per day (in contrast to the almost 9 h in the summer). A discussion about SOA formation mechanisms alternative to gas-phase photochemistry is presented later in section 4.2.

At the rural site of San Pietro Capofiume, as expected, the dominant contribution to POA in the cold periods is provided by
20 BBOA (varying between 28 and 33% of total OA mass during 2013 and 2011 fall campaigns, respectively) and the fraction of OOA to total OA is larger than at the urban site (35-65% in fall, and reaching 96% in summer). Peculiar results were found for the SPC fall 2011 campaign, during which very large contributions of POA were recorded: the HOA fraction reached 32% of OA mass, somewhat strange for a rural site. [Gilardoni et al. \(2014\) specifically studied this campaign suggesting that this high HOA relative contributions are](#) likely due to the occurrence of persistent fogs, [which are](#) scavenging
25 the most water-soluble OA components and leaving the interstitial aerosol enriched in its most hydrophobic organic components ([i.e.](#) HOA) (Gilardoni et al., 2014).

A summary of the seasonality of OA fractions at the two Po Valley sites is shown in Figure 5. The COA fraction, that was determined only at BO during one individual campaign and in small amounts, was not considered here to simplify the comparison between the other components. The SPC fall 2011 campaign was also not included in this statistic since the
30 aerosol composition and concentrations for this experiment referred to a mixture of total OA and interstitial OA in fog conditions, as [mentioned above and better](#) described [in Gilardoni et al. \(2014\)](#).

Table S10 reports the correlation coefficients between the PMF factors discussed so far and the main chemical species constituting the sub-micrometric aerosol masses measured by the HR-TOF-AMS. The highest correlations are observed

Marco Paglione 25/7/y 18:44

Eliminato: analyzed by itself

Marco Paglione 31/7/y 19:02

Eliminato: S

Marco Paglione 31/7/y 15:50

Eliminato: This was

Marco Paglione 1/8/y 15:07

Eliminato: above.

between OOA and secondary inorganic species, nitrate and ammonium sulfate, confirming the secondary nature of this fraction of OA. In particular, it can be noticed that OOA correlates better with ammonium nitrate in winter and fall, and with ammonium sulfate in summer and late spring, in agreement with previous results (Zhang et al., 2011). This behavior likely reflects the differences in temperature and relative humidity between winter and summer, which shift the partitioning of nitrate toward gas-phase (due to its volatility) during warm season. In addition, the different correlation suggests the possibility of a different oxidation pathways in secondary species formation between cold and warm seasons: a pathway characterized by cold temperature and high relative humidity (correlating with nitrate) and another one related with higher temperature and photochemical activity (correlating more with sulfate). The latter hypothesis will be better developed in the following sections.

For the campaigns carried out in parallel at the urban and rural site (summer 2012, spring 2013 and fall 2013), we estimated an “urban increment”, i.e., the increase in OA-type concentrations in urban areas with respect to the regional background. We expressed the increment as the ratio between the campaign average concentrations at the urban vs rural site, accordingly to season and the specific OA fraction considered (see Table 3). For total organic aerosol (OA) and for its OOA fraction, the ratios are quite constant throughout the seasons, varying between 1.13-1.36 and 0.97-1.30, respectively. By contrast, higher values were found for HOA (1.67, 1.91 and 2.85 in spring, fall and summer, respectively), in agreement with a major HOA source from urban traffic. The urban increment of BBOA is less clear: it varies a lot between spring (in which its value is very high, i.e. 5.87) and fall (with 0.67). Nevertheless, the spring value is affected by the low and intermittent high BBOA levels likely indicating very local sources. The fall value seems more representative and suggests a higher contribution of BBOA in the rural areas, probably due to the more spread use of fire-places and wood-stoves for domestic heating and to additional possible sources, such as agricultural burning.

4 SOA sources and their evolution

In the previous section we presented OOA as one single component; however, the HR-TOF-AMS statistical analysis identified various OOA types that may indicate different formation (sources) and transformation processes (aging) of SOA in the aerosol. The number of OOA categories identified during the SUPERSITO campaigns ranged from one (for SPC fall 2011 campaign) to four (for SPC summer 2012 campaign). Most of the IOPs (7 out of 12) allowed the identification of three OOA factors.

The spectral profiles of the individual OOAs are distinguishable based on minor mass fragments and other parameters. Among the most common parameters used in literature for the distinction and interpretation of the various OOA factors, are: elemental ratios (OM:OC, O:C and H:C), the carbon oxidation state (OSc) and the fractional abundance f (where $f\#$ is the ratio between the abundance of a specific ion and the total organic spectrum) of specific fragments in their spectral profiles

Marco Paglione 6/8/y 15:45
Formattato: Non Evidenziato

Marco Paglione 2/8/y 10:58
Eliminato: -

Marco Paglione 1/8/y 19:22
Formattato: Tipo di carattere:Non

Marco Paglione 26/8/y 10:23
Eliminato: -

Marco Paglione 31/7/y 13:45
Eliminato: ,

(e.g., CO_2^+ at m/z 44 (f_{44}); $\text{C}_2\text{H}_3\text{O}^+$ at m/z 43 (f_{43}); $\text{C}_2\text{H}_4\text{O}_2^+$ at m/z 60 (f_{60}); etc.). The elemental ratios and the relative proportion between f_{43} and f_{44} generally indicate the degree of oxidation and therefore the extent of aging of a single factor (normally the less oxidized components exhibit higher H:C, lower O:C and less f_{43} and f_{44} , while OOAs have O:C and f_{44} increasing with their degree of oxidation and aging in the atmosphere, largely due to the formation of carboxylic acids during this process) (Ng et al., 2010; Duplissy et al., 2011).

Tables S12 and S13 show a summary of the parameters for the analysis and interpretation of all the factors identified by the PMF statistical analysis (including the different OOAs listed in order of their O:C ratios) during the campaigns of the SUPERSITO project. We focus on two aspects: the influence of biomass burning emissions on OOA components and the importance of the aqueous-phase processing in their formation and evolution. A more comprehensive analysis of the OOAs features of particular IOPs is object of specific publications (Gilardoni et al., 2014; Sullivan et al., 2016; Gilardoni et al., 2016; Zanca et al., in preparation).

4.1 Biomass burning influence on SOA

The products of cellulose pyrolysis, such as levoglucosan and similar species (i.e., mannosan, galactosan, etc., collectively called hereinafter “anhydrosugars”), generate mass spectra with an enhanced signal at m/z 60 and 73 due to the ions $\text{C}_2\text{H}_4\text{O}_2^+$ and $\text{C}_3\text{H}_5\text{O}_2^+$, which are therefore considered good tracers of wood combustion (Schneider et al., 2006; Alfara et al., 2007). So, the parameter f_{60} (the ratio of the integrated signal at m/z 60 to the total signal of OA mass spectrum) is used as a marker to evaluate the influence of biomass burning emissions to the OA components (Cubison et al., 2011).

Fresh biomass burning emissions (BBOA factors) exhibit the highest content of anhydrosugars (f_{60}). During atmospheric aging, the relative intensity of anhydrosugars signal decreases because of degradation and oxidation reactions. At the same time, atmospheric aging leads to the oxidation of the molecules, which corresponds to the increase of oxygenated fragments in the mass spectrum, the most intense of which is at m/z 44 (CO_2^+ , f_{44}).

The contribution of f_{60} on the different OA components of each campaign is represented in Figure 6 by points in the f_{44} vs f_{60} space (Cubison et al., 2011) together with those of some references from previous studies (Aiken et al., 2009; Ng et al., 2011; Mohr et al., 2012; Saarikoski et al., 2012; Crippa et al., 2014; Florou et al., 2017). The background level indicating no influence of biomass burning is represented in Figure 6 (panels a -c) by a grey shaded area. As additional reference of OA not influenced by biomass combustion, we also report the measurements carried out during the summer 2012 parallel campaign at the high altitude background station of Mount Cimone (Rinaldi et al., 2015).

Figure 6 (panels a-c) shows that the spectral features of the OOA factors from several campaigns are those typical of aged OA (large f_{44}) but also indicate the presence of anhydrosugars above the background level. This suggests a variable influence of biomass combustion on the OOA factors.

Such OOAs factors influenced by biomass burning (OOAx_BB) represent a substantial mass fraction of the total OA during the fall-winter period (17-61% at the Bologna site and 14-35% at SPC). In the spring season, the biomass burning impact on

OOA composition is much less evident (f_{60} closer to the background levels) but still representing 37% of the total OA, more than twice the contribution of POA at BO during the spring 2013 campaign.

Additional tests and details on the determination of the biomass burning influence on OOA components are discussed in the supplemental section S2.2.3.

5

4.2 Biomass burning oxidation pathways

The vertical axis in Figure 6 is controlled by the oxidation of the bulk OA while the horizontal axis by the anhydrosugars loss. Thus, depending on the relative rates of these processes, the slopes of the virtual lines connecting the primary factors (BBOAs) and the corresponding aged PMF factors (OOAx_BB) are expected to be different. We do see indeed that slopes vary in different campaigns. We also see that two OOA_BB factors detected during BO fall 2011 and winter 2013 campaigns are connected to the primary BBOA with different slopes in the f_{60} vs f_{44} space (as reported by the arrows in Figure 6a). This variability could suggest that the two OOA_BB components, observed during the same experiment, are formed through different oxidation rates and pathways due to the variable environmental conditions.

15 In order to test this hypothesis, the evolution of the BBOA into OOAs is further analyzed for BO fall 2011 and BO winter 2013 in Figure 6d using the O:C and the hydrogen-to-carbon (H:C) ratios of the BBOA and OOAx_BB factors in the Van Krevelen (VK) diagram. The VK diagram is typically used to investigate the OA evolution during field and laboratory experiments (Heald et al., 2010; Ng et al., 2011). The plot allows to remove the effect of physical mixing between secondary and primary aerosols, providing a clearer interpretation of the results. Aerosol aging has the overall effect of increasing O:C ratios. In the VK plot the H:C vs. O:C slope of 0 is equivalent to the replacement of a hydrogen atom with an OH moiety, whereas a slope of -1 indicates the formation of carboxylic acid groups (Ng et al., 2011). O:C and H:C values are reported for BBOA (triangles), OOA_BB factors (squares and circles). The slope of the line that links BBOA to the circles (i.e., OOAx_BB-aq) is close to zero while the line linking BBOA to the squares (i.e., OOAx_BB) is between -0.5 and -1, suggesting possible different oxidation pathways. The negative slope indicates that OOAx_BB likely formed from BBOA through formation of carboxylic acid moieties, suggesting photochemical oxidation processes driven by OH radical which might take place both on gas and aqueous phase (Ng et al., 2011; Timonen et al., 2013; McNeill, 2015). Conversely, OOAx_BB-aq formation (slope 0) is consistent with the hydroxyl group formation possibly taking place in aerosol water (i.e., wet aerosol) through dark chemistry (Lim et al., 2010; Gilardoni et al., 2016), an hypothesis better assessed in the next paragraph.

30 4.3 Aqueous-phase chemistry in SOA formation

Figure 7 shows the variations in contributions of the two BB-influenced OOA factors identified during the BO fall2011 campaign as a function of RH, together with some other meteorological and chemical parameters. The aerosol liquid water

Marco Paglione 16/7/y 18:37

Eliminato: squares

Marco Paglione 16/7/y 18:37

Eliminato: circles

Marco Paglione 16/7/y 18:26

Eliminato: -aq

Marco Paglione 25/7/y 17:08

Eliminato: took place in aerosol water (i.e., wet aerosol) through dark chemistry

content (ALWC), as calculated by the ISORROPIA-II model, and the hydroxymethanesulfonate (HMSA) were used to trace the effects of aqueous-phase SOA formation. In fact HMSA is formed by the reaction of sulfite and bisulfite with dissolved formaldehyde in droplets and deliquesced aerosols and is oxidized by ozone at concentration as low as 10 ppb (Kok et al., 1986; Facchini et al., 1992; Whiteaker and Prather, 2003). We suggest here that formaldehyde (as well as a number of other gaseous compounds including ketones, aldehydes, and small carboxylic acids in the BB plumes, Schauer et al., 2001; Andreae, 2019) would preferentially partition into particles at high ALWC and would react to form HMSA (and/or other products). Then the products of these aqueous phase reactions (such as HMSA) remain in the particle phase after water evaporation, changing the chemical composition of the organic aerosol. For this reason, considering also that HMSA formation is inhibited by photochemistry (due to its fast reaction with ozone) and that the analyzed aerosol was dried before sampling, the correlation of some factors with HMSA can be considered a reliable evidence of aqueous-phase formation pathway of some OA fractions. HMSA was detected by the HR-TOF-AMS (following the estimation method presented by Ge et al., 2012) during all the campaigns, and its presence and concentration was confirmed by off-line H-NMR analysis of filter samples. ALWC and HMSA exhibit a strong increase as a function of RH during the campaign confirming the influence of aqueous-phase processing at high RH levels (Figure 7, panel a). At the same time, temperature and solar radiation (Figure 7, panel b) decrease as function of RH suggesting a reduction of photochemical activity.

These ambient conditions result into a large increase in the contribution of OOA2_BB-aq, whereas the OOA1_BB concentration remained relatively constant (Figure 7c).

Extending this analysis to all the campaigns (see also Figure S5), we identified at least one OOA factor originating from biomass burning through aqueous-phase processing (OOAx_BB-aq) in 8 out of 12 datasets (all fall and winter campaigns plus spring 2013). The correlations of all the OOAs with the aerosol liquid water content (ALWC) and the hydroxymethanesulfonate (HMSA) are summarized in Table 4 (and reported also in Figure 8 for the OOAx_BB-aq factors).

The spectral profiles of these OOA_BB-aq factors originated from aqueous-phase processing (reported in Figure 8) are characterized by higher signals at m/z 29 (CHO^+) and m/z 58 ($\text{C}_2\text{H}_2\text{O}_2^+$) in addition to the more common m/z 43 ($\text{C}_2\text{H}_3\text{O}^+$), m/z 44 (CO_2^+) and m/z 60 ($\text{C}_2\text{H}_4\text{O}_2^+$) that characterized also the other BB-influenced secondary components. The OOAx_BB-aq factors spectra have also good similarities ($4 < \theta \text{ angle} < 29$, see Table S15) between each other and with the OOA spectra recorded after fog dissipation at SPC during fall 2011 (Gilardoni et al., 2016).

The conclusion that these components are affected by aqueous-phase processing is further supported by the correlations between the OOAx_BB-aq factors and some specific fragment ions. As shown in Table S16 all the aqSOAs identified during SUPERSITO campaigns are well correlated with $\text{C}_2\text{H}_2\text{O}_2^+$, C_2O_2^+ and CH_2O_2^+ , which are typical fragments of methylglyoxal and glyoxal, that are precursors of SOA via cloud processing (Carlton et al., 2007; Altieri et al., 2008). We further stress the link between biomass burning and these aqSOA by looking at the correlations of these components with specific fragment ions of aqueous-phase products of phenol and guaiacol emitted during the biomass burning (namely PhOH-OH, $\text{C}_6\text{H}_6\text{O}_2^+$, m/z 110.037; PhOH-2OH, $\text{C}_6\text{H}_6\text{O}_3^+$ at m/z 126.032; GUA-OH, $\text{C}_7\text{H}_8\text{O}_3^+$ at m/z 140.047; GUA-2OH, $\text{C}_7\text{H}_8\text{O}_4^+$ at m/z

Marco Paglione 6/8/y 15:45
Formattato: Non Evidenziato

Marco Paglione 6/8/y 15:45
Formattato: Non Evidenziato

Marco Paglione 6/8/y 15:45
Formattato: Non Evidenziato

Marco Paglione 30/7/y 14:13
Eliminato: indicating the possible

Marco Paglione 30/7/y 14:13
Eliminato: s

Marco Paglione 25/7/y 19:06
Eliminato: Dividing the individual OOA fractions for the total POA concentrations (considered as a surrogate of the planetary boundary layer, PBL) in Figure 7d, we observed similar variations of the BB-SOA components with $\text{RH} > 60\%$, further supporting the above conclusion.

Marco Paglione 30/7/y 14:31
Eliminato: -

156.042), already identified in previous laboratory studies (Yu et al., 2014). Moreover, considering the elemental composition of the OOAx_BB-aq (Table S12-S13 and Figure 8) we notice that their O:C ratios, calculated following the Ambient Improved (AI) method (Canagaratna et al., 2015), are similar (on average 0.82 ± 0.09) to the AI O:C ratios obtained from laboratory oxidation of phenolic compounds (0.89 ± 0.10 , Sun et al., 2010; 1.03 ± 0.17 , Yu et al., 2014) and from laboratory-generated SOA from the photooxidation of organic precursors in the aqueous phase (0.89 ± 0.13 , Lee et al., 2011; Lee et al., 2012).

In conclusion, BB-influenced SOA formed by aqueous-phase processing (bb-aqSOA) identified during the SUPERSITO campaigns represents a substantial mass fraction of the total OA during fall-winter months (14-28% at Bologna site and 14-35% at SPC). This component is often more than half of the total SOA influenced by BB-emissions, while the other half undergoes photochemical oxidation pathways leading to OOAx_BB. Overall, our results support the importance in the Po Valley of SOA formation by aqueous-phase processing of wood combustion reported by Gilardoni et al. (2016), extending the ambient observations of these phenomena to a larger dataset (Figure 9).

15 5 Conclusions

The SUPERSITO project constitutes the first extensive (multi-sites and multi-years) time-resolved aerosol chemical experiment in the Po Valley. Eight intensive observation periods (IOPs) were carried out over the four years of the project (from 2011 to 2014) at two different sites (Bologna, urban background and San Pietro Capofiume, rural background) using a High Resolution Aerosol Mass Spectrometer (HR-AMS). The source apportionment of the OA allowed improving our understanding of aerosol sources, their chemical features and spatial-temporal variability in the region, one of the most important pollution hot spots in Europe. Considering the special focus of the project on the cold season (3 campaigns in fall and 2 in winter, out of 8 in total) it was especially possible to investigate the wintertime SOA formation pathways, which are the less characterized and, for this reason, one of the most important missing processes in atmospheric chemistry and air quality models (Tsimpidi et al., 2016).

The possibility to compare the organic factors identified by the HR-AMS with additional chemical tracers measured in parallel by other advanced spectroscopic techniques (i.e., NMR) and more traditional ones (e.g., IC, GC/MS, OC/EC, etc.) provided new insights on the detailed chemical structure and especially on the formation and ageing mechanisms of SOA.

On the multi-years basis of the project, OA represent on average $45 \pm 8\%$ (33-58%) and $46 \pm 7\%$ (36-50%) of the total non-refractory submicron particles (PM1-NR) at the urban and rural site respectively, within the range reported in literature for other European sites (Crippa et al., 2014) and the Asian regions (Hu et al., 2017; Li et al., 2015; Wu et al., 2018 for China & East Asia; Chakraborty et al., 2018 for India), and slightly less than the values reported for Southeastern US (50-75%, Xu et al., 2015; Budisulistiorini et al., 2016). Among this fraction, primary sources (POA) are dominated by biomass burning

(23±13%), especially at the rural site (SPC) whereas the fossil fuel combustion (12±7%) is higher in the urban background site (Bologna) where it also presents a marked seasonality. However, the biomass burning contribution to POA remains the most important source of POA also at the urban site during the cold fall/winter seasons. The BBOA contribution ranging 17-38% at Bologna during the fall/winter seasons is not far from the values reported for other European cities (10-40% in Paris, Crippa et al., 2013b; 5-27% from the EUCAARI multi-sites study, Crippa et al., 2014) and United States areas (e.g., 15-33% for Southeast US, Budisulistiorini et al., 2016) and slightly higher than that of other highly populated and polluted cities/regions of Asia (11-14% at Beijing, China, Sun et al., 2018; 10-20% at Kunpur, India, Chakraborty et al., 2018) where, however, other combustion sources (i.e., coal) contribute to the POA fraction.

The contribution of oxidized organic aerosol (OOA, used as a proxy for SOA) were found to be much higher than the primary ones, regardless of site and season with a multi-year average of 66% (44-92%; st.dev.=16%) and 71% (35-96%; st.dev.=27%) of the total OA mass, at the urban and rural site respectively. The SOA dominance is also observed during winter at the urban site, where the SOA represents on average 56% (50-61%; st.dev.=8%) of the total OA mass. Within this SOA, the measurements highlight the dominant presence of biomass burning secondary components, even in the urban background. The HR-AMS data indicate that the OA mass contributions of this SOA factor influenced by wood-combustion was of the order of 14-44% which translates into biomass burning emissions representing the 31-82% of the OM mass in the Po Valley during cold months (fall and winter). Significant contribution of aged BB emissions on the OA mass loadings has been already suggested by previous studies regarding the Po Valley (Saarikoski et al., 2012) and different European (Paris, France, Crippa et al., 2013b; Eastern Mediterranean, Bougiatioti et al., 2014; Athens, Greece, Stavroulas et al., 2018), Asian (Beijing, China, Hu et al., 2017 & Sun et al., 2018; Kunpur, India, Chakraborty et al., 2018) and American sites (South-East US, Xu et al., 2015; Budisulistiorini et al., 2016). However, studies reporting the identification and quantification in ambient air of specific BB-influenced OOA factors are still very limited (Gilardoni et al., 2016; Xu et al., 2017).

Our study also identified and quantified a particularly relevant role of the aqueous-phase processing in the formation and transformation of primary biomass burning emissions. Aqueous SOA (aqSOA) factors identified as OOAx_BB-aq represent on average 21% (14-28%) and 25% (14-35%) of the total OA mass at the urban and rural sites, respectively, highlighting the importance of aqueous-phase processing for SOA formation and transformation. Considering the widespread wintertime occurrence of fog, low-level clouds and wet aerosols in many other highly-populated sites enclosed in orographic basins (Benelux and Ruhr district, Paris and London basins, Cermak et al., 2009; Californian Central Valley, Baldocchi et al., 2014; Yangtze River corridor, Niu et al., 2010; and Indo-Gangetic plain, Saraf et al., 2011), this study strongly suggests that aqueous-processing can be a major driver for secondary aerosol formation in wintertime at all these sites, with important consequences on air quality policy at the global level.

These results suggest the importance of a continuous monitoring system for better characterization of biomass burning-driven pollution in the Po Valley area, using complementary measurements both routinely and through intensive campaigns in order to explore the importance of biomass burning on air quality and climate.

Author contributions. M.P., S.G., S.D. S.F., and M.C.F. designed the research; V.P., S.G., S.D., and M.C.F. organized the field campaigns; M.P., S.G., M.R., L.P., P.M., and C.C. carried out the experiments and processed AMS data; M.P., S.G., L.P., and P.M. performed the AMS PMF; L.P., F.C., A.S.H.P., contributed to the PMF discussion and correction; M.P., S.G., N.Z., S.S., L.G., D.B., S.F., F.S., and A.T. collected and analyzed the aerosol samples. M.P., S.G., M.R., and S.D. wrote the paper. H.H., A.W., and S.F. contributed the scientific discussion and paper correction.

Acknowledgments. This work is supported by Emilia-Romagna Region “Supersito” Project (DRG 428/10; DGR 1971/2013) and European projects PEGASOS (EU FP7-ENV-2010-265148) and BACCHUS (EU FP7-ENV-2013-603445). Authors acknowledge COST Action CA16109 COLOSSAL. Authors are grateful to all the ARPAE Emilia-Romagna co-workers for their support and collaboration in the realization of the field campaigns. ISAC-CNR is particularly grateful to Leone Tarozzi, Francescopiero Calzolari and all the other colleagues who collaborated in the preparation of the campaigns, the functioning of the instruments and the aerosol sampling in the field.

References

Aiken, A. C., Salcedo, D., Cubison, M. J., Huffman, J. A., DeCarlo, P. F., Ulbrich, I. M., Docherty, K. S., Sueper, D., Kimmel, J. R., Worsnop, D. R., Trimborn, A., Northway, M., Stone, E. A., Schauer, J. J., Volkamer, R. M., Fortner, E., de Foy, B., Wang, J., Laskin, A., Shutthanandan, V., Zheng, J., Zhang, R., Gaffney, J., Marley, N. A., Paredes-Miranda, G., Arnott, W. P., Molina, L. T., Sosa, G., and Jimenez, J. L.: Mexico City aerosol analysis during MILAGRO using high resolution aerosol mass spectrometry at the urban supersite (T0) - Part 1: Fine particle composition and organic source apportionment, *Atmos. Chem. and Phys.*, 9, 6633–6653, doi:10.5194/acp-9-6633-2009, 2009.

Alfarra, M. R., Prevot, A. S. H., Szidat, S., Sandradewi, J., Weimer, S., Lanz, V. A., Schreiber, D., Mohr, M., and Baltensperger, U.: Identification of the mass spectral signature of organic aerosols from wood burning emissions, *Environ. Sci. Technol.*, 41, 5770–5777, 2007.

Altieri, K. E., Seitzinger, S. P., Carlton, A. G., Turpin, B. J., Klein, G. C.; Marshall, A. G.: Oligomers formed through in-cloud methylglyoxal reactions: Chemical composition, properties, and mechanisms investigated by ultra-high resolution FT-ICR mass spectrometry. *Atmos. Environ.*, 42 (7), 1476–1490, 2008.

[Andreae, M. O.: Emission of trace gases and aerosols from biomass burning – an updated assessment, *Atmos. Chem. Phys.*, 19, 8523-8546, https://doi.org/10.5194/acp-19-8523-2019, 2019.](https://doi.org/10.5194/acp-19-8523-2019)

Baldocchi, D., and Waller, R.: Winter fog is decreasing in the fruit growing region of the Central Valley of California, *Geophys. Res. Lett.*, 41, 3251–3256, 2014.

- Belis, C. A., Karagulian, F., Larsen, B. R., and Hopke, P. K.: Critical review and meta-analysis of ambient particulate matter source apportionment using receptor models in Europe, *Atmos. Environ.*, 69, 94–108, doi:10.1016/j.atmosenv.2012.11.009, 2013.
- 5 Bougiatioti, A., Stavroulas, I., Kostenidou, E., Zampas, P., Theodosi, C., Kouvarakis, G., Canonaco, F., Prévôt, A. S. H., Nenes, A., Pandis, S. N., and Mihalopoulos, N.: Processing of biomass-burning aerosol in the eastern Mediterranean during summertime, *Atmos. Chem. Phys.*, 14, 4793–4807, <https://doi.org/10.5194/acp-14-4793-2014>, 2014.
- Bressi, M., Sciare, J., Ghersi, V., Bonnaire, N., Nicolas, J. B., Petit, J.-E., Moukhtar, S., Rosso, A., Mihalopoulos, N., and
10 Féron, A.: A one-year comprehensive chemical characterisation of fine aerosol (PM_{2.5}) at urban, suburban and rural background sites in the region of Paris (France), *Atmos. Chem. Phys.*, 13, 7825–7844, <https://doi.org/10.5194/acp-13-7825-2013>, 2013.
- Bressi, M., Cavalli, F., Belis, C. A., Putaud, J.-P., Fröhlich, R., Martins dos Santos, S., Petralia, E., Prévôt, A. S. H., Berico,
15 M., Malaguti, A., and Canonaco, F.: Variations in the chemical composition of the submicron aerosol and in the sources of the organic fraction at a regional background site of the Po Valley (Italy), *Atmos. Chem. Phys.*, 16, 12875–12896, <https://doi.org/10.5194/acp-16-12875-2016>, 2016.
- Budisulistiorini, S. H., Baumann, K., Edgerton, E. S., Bairai, S. T., Mueller, S., Shaw, S. L., Knipping, E. M., Gold, A., and
20 Surratt, J. D.: Seasonal characterization of submicron aerosol chemical composition and organic aerosol sources in the southeastern United States: Atlanta, Georgia, and Look Rock, Tennessee, *Atmos. Chem. Phys.*, 16, 5171–5189, <https://doi.org/10.5194/acp-16-5171-2016>, 2016.
- Canagaratna, M. R., Jayne, J. T., Ghertner, D. A., Herndon, S., Shi, Q., Jimenez, J. L., Silva, P. J., Williams, P., Lanni, T.,
25 Drewnick, F., Demerjian, K. L., Kolb, C. E., and Worsnop, D. R.: Chase studies of particulate emissions from in-use New York City vehicles, *Aerosol Sci. Technol.*, 38, 555–573, 2004.
- Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q., Onasch, T. B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R., Trimborn, A. M., Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits,
30 P., and Worsnop, D. R.: Chemical and microphysical characterization of ambient aerosols with the Aerodyne aerosol mass spectrometer, *Mass Spectrom. Rev.*, 26, 185–222, doi:10.1002/mas.20115, 2007.
- Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P., Hildebrandt Ruiz, L., Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue, N. M., Jayne, J. T., and Worsnop, D. R.: Elemental ratio

- measurements of organic compounds using aerosol mass spectrometry: characterization, improved calibration, and implications, *Atmos. Chem. Phys.*, 15, 253–272, doi:10.5194/acp-15-253-2015, 2015.
- 5 Canonaco, F., Crippa, M., Slowik, J. G., Baltensperger, U., and Prévôt, A. S. H.: SoFi, an IGOR-based interface for the efficient use of the generalized multilinear engine (ME-2) for the source apportionment: ME-2 application to aerosol mass spectrometer data, *Atmos. Meas. Tech.*, 6, 3649–3661, doi:10.5194/amt-6-3649-2013, 2013.
- 10 Canonaco, F., Slowik, J. G., Baltensperger, U., and Prévôt, A. S. H.: Seasonal differences in oxygenated organic aerosol composition: implications for emissions sources and factor analysis, *Atmos. Chem. Phys.*, 15, 6993–7002, doi:10.5194/acp-15-6993-2015, 2015.
- 15 Carbone, C., Decesari, S., Paglione, M., Giulianelli, L., Rinaldi, M., Marinoni, A., Cristofanelli, P., Diodato, A., Bonasoni, P., Fuzzi, S., and Facchini, M. C.: 3-year chemical composition of free tropospheric PM₁ at the Mt. Cimone GAW global station – South Europe – 2165m a.s.l., *Atmos. Environ.*, 87, 218–227, doi:10.1016/j.atmosenv.2014.01.048, 2014.
- 20 Carlton, A. G., Turpin, B. J., Altieri, K. E., Seitzinger, S., Reff, A., Lim, H.-J.; Ervens, B.: Atmospheric oxalic acid and SOA production from glyoxal: Results of aqueous photooxidation experiments. *Atmos. Environ.*, 41 (35), 7588–7602, 2007.
- Cavalli, F., Viana, M., Yttri, K. E., Genberg, J., and Putaud, J.-P.: Toward a standardised thermal-optical protocol for measuring atmospheric organic and elemental carbon: the EUSAAR protocol, *Atmos. Meas. Tech.*, 3, 79–89, <https://doi.org/10.5194/amt-3-79-2010>, 2010.
- 25 Cermak, J., Eastman, R. M., Bendix, J., and Warren S. G.: European climatology of fog and low stratus based on geostationary satellite observations, *Q. J. R. Meteorol. Soc.*, 135, 2125–2130, 2009.
- Chakraborty A., Mandariya A. K., Chakraborti R., Gupta T., Tripathi S.N.: Realtime chemical characterization of post monsoon organic aerosols in a polluted urban city: Sources, composition, and comparison with other seasons, *Environ Pollut.*, 232 : 310–321. doi: 10.1016/j.envpol.2017.09.079, 2018.
- 30 Chen Q., Liu Y.J., Donahue N.M., Shilling J.E., Martin S.T.: Particle-phase chemistry of secondary organic material: modeled compared to measured O: C and H: C elemental ratios provide constraints, *Environ. Sci. Technol.*, 45, p. 4763, 2011.

- Costabile, F., Gilardoni, S., Barnaba, F., Di Ianni, A., Di Liberto, L., Dionisi, D., Manigrasso, M., Paglione, M., Poluzzi, V., Rinaldi, M., Facchini, M. C., and Gobbi, G. P.: Characteristics of brown carbon in the urban Po Valley atmosphere, *Atmos. Chem. Phys.*, 17, 313-326, <https://doi.org/10.5194/acp-17-313-2017>, 2017.
- 5 Crippa, M., Haddad, I. E., Slowik, J. G., Decarlo, P. F., Mohr, C., Heringa, M. F., Chirico, R., Marchand, N., Sciare, J., Baltensperger, U., and Prevot, A. S. H.: Identification of marine and continental aerosol sources in Paris using high resolution aerosol mass spectrometry, *J. Geophys. Res.-Atmos.*, 118, 1950–1963, <https://doi.org/10.1002/jgrd.50151>, 2013a.
- Crippa, M., DeCarlo, P. F., Slowik, J. G., Mohr, C., Heringa, M. F., Chirico, R., Poulain, L., Freutel, F., Sciare, J., Cozic, J.,
10 Di Marco, C. F., Elsasser, M., José, N., Marchand, N., Abidi, E., Wiedensohler, A., Drewnick, F., Schneider, J., Borrmann, S., Nemitz, E., Zimmermann, R., Jaffrezo, J.-L., Prévôt, A. S. H., and Baltensperger, U.: Wintertime aerosol chemical composition and source apportionment of the organic fraction in the metropolitan area of Paris, *Atmos. Chem. Phys.*, 13, 961–981, [doi:10.5194/acp-13-961-2013](https://doi.org/10.5194/acp-13-961-2013), 2013b.
- 15 Crippa, M., Canonaco, F., Lanz, V. A., Äijälä, M., Allan, J. D., Carbone, S., Capes, G., Ceburnis, D., Dall'Osto, M., Day, D. A., DeCarlo, P. F., Ehn, M., Eriksson, A., Freney, E., Hildebrandt Ruiz, L., Hillamo, R., Jimenez, J. L., Junninen, H., Kiendler-Scharr, A., Kortelainen, A.-M., Kulmala, M., Laaksonen, A., Mensah, A. A., Mohr, C., Nemitz, E., O'Dowd, C., Ovadnevaite, J., Pandis, S. N., Petäjä, T., Poulain, L., Saarikoski, S., Sellegri, K., Swietlicki, E., Tiitta, P., Worsnop, D. R., Baltensperger, U., and Prévôt, A. S. H.: Organic aerosol components derived from 25 AMS data sets across Europe using a
20 consistent ME-2 based source apportionment approach, *Atmos. Chem. Phys.*, 14, 6159-6176, <https://doi.org/10.5194/acp-14-6159-2014>, 2014.
- Cubison, M. J., Ortega, A. M., Hayes, P. L., Farmer, D. K., Day, D., Lechner, M. J., Brune, W. H., Apel, E., Diskin, G. S., Fisher, J. A., Fuelberg, H. E., Hecobian, A., Knapp, D. J., Mikoviny, T., Riemer, D., Sachse, G. W., Sessions, W., Weber, R.
25 J., Weinheimer, A. J., Wisthaler, A., and Jimenez, J. L.: Effects of aging on organic aerosol from open biomass burning smoke in aircraft and laboratory studies, *Atmos. Chem. Phys.*, 11, 12049–12064, [doi:10.5194/acp-11-12049-2011](https://doi.org/10.5194/acp-11-12049-2011), 2011.
- Dall'Osto, M., Paglione, M., Decesari, S., Facchini, M. C., O'Dowd, C., Plass-Duellmer, C., and Harrison, R. M.: On the Origin of AMS “Cooking Organic Aerosol” at a Rural Site, *Environ. Sci. Technol.*, 49, 13964–13972,
30 [doi:10.1021/acs.est.5b02922](https://doi.org/10.1021/acs.est.5b02922), 2015.
- DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C., Gonin, M., Fuhrer, K., Horvath, T., Docherty, K. S., Worsnop, D. R., and Jimenez, J. L.: Field-Deployable, High-Resolution, Time-of-Flight Aerosol Mass Spectrometer, *Anal. Chem.*, 78, 8281–8289, <https://doi.org/10.1021/ac061249n>, 2006.

- 5 DeCarlo, P. F., Dunlea, E. J., Kimmel, J. R., Aiken, A. C., Sueper, D., Crounse, J., Wennberg, P. O., Emmons, L., Shinozuka, Y., Clarke, A., Zhou, J., Tomlinson, J., Collins, D. R., Knapp, D., Weinheimer, A. J., Montzka, D. D., Campos, T., and Jimenez, J. L.: Fast airborne aerosol size and chemistry measurements above Mexico City and Central Mexico during the MILAGRO campaign, *Atmos. Chem. Phys.*, 8, 4027–4048, doi:10.5194/acp-8-4027-2008, 2008.
- 10 Decesari, S., Fuzzi, S., Facchini, M. C., Mircea, M., Emblico, L., Cavalli, F., Maenhaut, W., Chi, X., Schkolnik, G., Falkovich, A., Rudich, Y., Claeys, M., Pashynska, V., Vas, G., Kourtchev, I., Vermeylen, R., Hoffer, A., Andreae, M. O., Tagliavini, E., Moretti, F., and Artaxo, P.: Characterization of the organic composition of aerosols from Rondônia, Brazil, during the LBASMOCC 2002 experiment and its representation through model compounds, *Atmos. Chem. Phys.*, 6, 375–402, doi:10.5194/acp-6-375-2006, 2006.
- 15 Decesari, S., Allan, J., Plass-Duelmer, C., Williams, B. J., Paglione, M., Facchini, M. C., O’Dowd, C., Harrison, R. M., Gietl, J. K., Coe, H., Giulianelli, L., Gobbi, G. P., Lanconelli, C., Carbone, C., Worsnop, D., Lambe, A. T., Ahern, A. T., Moretti, F., Tagliavini, E., Elste, T., Gilge, S., Zhang, Y., and Dall’Osto, M.: Measurements of the aerosol chemical composition and mixing state in the Po Valley using multiple spectroscopic techniques, *Atmos. Chem. Phys.*, 14, 12109–12132, doi:10.5194/acp-14-12109-2014, 2014.
- 20 Duplissy, J., DeCarlo, P.F., Dommen, J., Alfarra, M.R., Metzger, A., Barmapadimos, I., Prevot, A. S. H., Weingartner, E., Tritscher, T., Gysel, M., Aiken, A.C., Jimenez, J.L., Canagaratna, M.R., Worsnop, D.R., Collins, D.R., Tomlinson, J., and Baltensperger, U.: Relating hygroscopicity and composition of organic aerosol particulate matter, *Atmos. Chem. Phys.*, 11, 1155–1165, 2011.
- 25 EEA, Air Quality in Europe d 2016 Report. EEA Report No 28/2016 (<https://www.eea.europa.eu/publications/air-quality-in-europe-2016>), 2016.
- Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in cloud droplets and aqueous particles (aqSOA): a review of laboratory, field and model studies, *Atmos. Chem. Phys.*, 11, 11069–11102, <https://doi.org/10.5194/acp-11-11069-2011>, 2011.
- 30 Facchini, M. C., Fuzzi, S., Kessel, M., Wobrock, W., Jaeschke, W., Arends, B. G., Mols, J. J., Berner, A., Solly, I., Krusiz, C., Reischl, G., Pahl, S., Hallberg, A., Ogren, J. A., Fierlinger-Oberlinninger, H., Marzorati, A., and Schell, D.: The chemistry of sulfur and nitrogen species in a fog system. A multiphase approach, *Tellus 44B*, 505–521, 1992.

- Florou, K., Papanastasiou, D. K., Pikridas, M., Kaltsonoudis, C., Louvaris, E., Gkatzelis, G. I., Patoulias, D., Mihalopoulos, N., and Pandis, S. N.: The contribution of wood burning and other pollution sources to wintertime organic aerosol levels in two Greek cities, *Atmos. Chem. Phys.*, 17, 3145-3163, <https://doi.org/10.5194/acp-17-3145-2017>, 2017.
- 5 Fountoukis, C. and Nenes, A.: ISORROPIA II: a computationally efficient thermodynamic equilibrium model for K^+ - Ca^{2+} - Mg^{2+} - NH_4^+ - Na^+ - SO_4^{2-} - NO_3^- - Cl^- - H_2O aerosols, *Atmos. Chem. Phys.*, 7, 4639-4659, <https://doi.org/10.5194/acp-7-4639-2007>, 2007.
- Ge, X., Zhang, Q., Sun, Y., Ruehl, C. R., Setyan, A.: Effect of aqueous-phase processing on aerosol chemistry and size
10 distributions in Fresno, California, during wintertime. *Environ Chem* 9(3):221–235, 2012.
- Gilardoni, S., Vignati, E., Cavalli, F., Putaud, J. P., Larsen, B. R., Karl, M., Stenström, K., Genberg, J., Henne, S., and Dentener, F.: Better constraints on sources of carbonaceous aerosols using a combined ^{14}C – macro tracer analysis in a European rural background site, *Atmos. Chem. Phys.*, 11, 5685–5700, doi:10.5194/acp-11-5685-2011, 2011.
- 15 Gilardoni, S., Massoli, P., Giulianelli, L., Rinaldi, M., Paglione, M., Pollini, F., Lanconelli, C., Poluzzi, V., Carbone, S., Hillamo, R., Russell, L. M., Facchini, M. C., and Fuzzi, S.: Fog scavenging of organic and inorganic aerosol in the Po Valley, *Atmos. Chem. Phys.*, 14, 6967–6981, doi:10.5194/acp-14-6967-2014, 2014.
- 20 Gilardoni, S., Massoli, P., Paglione, M., Giulianelli, L., Carbone, C., Rinaldi, M., Decesari, S., Sandrini, S., Costabile, F., Gobbi, G. P., Pietrogrande, M. C., Visentin, M., Scotto, F., Fuzzi, S., and Facchini, M. C.: Direct observation of aqueous secondary organic aerosol from biomass burning emissions, *P. Natl. Acad. Sci. USA*, 113, 10013–10018, 2016.
- Giulianelli, L., Gilardoni, S., Tarozzi, L., Rinaldi, M., Decesari, S., Carbone, C., Facchini, M.C., Fuzzi, S.: Fog occurrence
25 and chemical composition in the Po valley over the last twenty years. *Atmos. Environ.* 98, 394e401, 2014.
- Grieshop, A. P., Donahue, N. M., and Robinson, A. L.: Laboratory investigation of photochemical oxidation of organic aerosol from wood fires 2: analysis of aerosol mass spectrometer data, *Atmos. Chem. Phys.*, 9, 2227–2240, doi:10.5194/acp-9-2227-2009, 2009.
- 30 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, Th. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J.

- D., Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic aerosol: current and emerging issues, *Atmos. Chem. Phys.*, 9, 5155-5236, <https://doi.org/10.5194/acp-9-5155-2009>, 2009.
- Heald, C. L., Kroll, J. H., Jimenez, J. L., Docherty, K. S., DeCarlo, P. F., Aiken, A. C., Chen, Q., Martin, S. T., Farmer, D. K., and Artaxo, P.: A simplified description of the evolution of organic aerosol composition in the atmosphere, *Geophys. Res. Lett.*, 37, L08803, <https://doi.org/10.1029/2010gl042737>, 2010
- Heringa, M. F., DeCarlo, P. F., Chirico, R., Tritscher, T., Dommen, J., Weingartner, E., Richter, R., Wehrle, G., Prévôt, A. S. H., and Baltensperger, U.: Investigations of primary and secondary particulate matter of different wood combustion appliances with a high-resolution time-of-flight aerosol mass spectrometer, *Atmos. Chem. Phys.*, 11, 5945–5957, [doi:10.5194/acp-11-5945-2011](https://doi.org/10.5194/acp-11-5945-2011), 2011.
- Hu, W., Hu, M., Hu, W.-W., Zheng, J., Chen, C., Wu, Y., and Guo, S.: Seasonal variations in high time-resolved chemical compositions, sources, and evolution of atmospheric submicron aerosols in the megacity Beijing, *Atmos. Chem. Phys.*, 17, 9979-10000, <https://doi.org/10.5194/acp-17-9979-2017>, 2017.
- IPCC: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, Chapter 7: Clouds and Aerosols, edited by: Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M., Cambridge University Press, Cambridge, UK and New York, NY, USA, 2013.
- Jayne, J. T., Leard, D. C., Zhang, X., Davidovits, P., Smith, K. A., Kolb, C. E., and Worsnop, D. R.: Development of an aerosol mass spectrometer for size and composition analysis of submicron particles, *Aerosol Sci. Technol.*, 33, 49–70, 2000.
- Jimenez, J. L., Jayne, J. T., Shi, Q., Kolb, C. E., Worsnop, D. R., Yourshaw, I., Seinfeld, J. H., Flagan, R. C., Zhang, X. F., Smith, K. A., Morris, J. W., and Davidovits, P.: Ambient aerosol sampling using the aerodyne aerosol mass spectrometer, *J. Geophys. Res.-Atmos.*, 108, 447–457, <https://doi.org/10.1029/2001JD001213>, 2003.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimojo, A., Sun, J. Y., Zhang, Y. M.,

- Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of organic aerosols in the atmosphere, *Science*, 326, 1525–1529, doi:10.1126/science.1180353, 2009.
- 5 Kleinman, L. I., Springston, S. R., Daum, P. H., Lee, Y.-N., Nunnermacker, L. J., Senum, G. I., Wang, J., Weinstein-Lloyd, J., Alexander, M. L., Hubbe, J., Ortega, J., Canagaratna, M. R., and Jayne, J.: The time evolution of aerosol composition over the Mexico City plateau, *Atmos. Chem. Phys.*, 8, 1559–1575, <https://doi.org/10.5194/acp-8-1559-2008>, 2008.
- Kok G. L., Gitlin S. N., and Lazrus A. L.: Kinetics of the formation and decomposition of hydroxymethanesulfonate, *J. Geophys. Res.-Atmos.*, 91(D2):2801–2804, 1986.
- 10
- Kostenidou, E., Lee, B.-H., Engelhart, G. J., Pierce, J. R., and Pandis, S. N.: Mass spectra deconvolution of low, medium and high volatility biogenic Secondary Organic Aerosol, *Environ. Sci. Technol.*, 43, 4884–4889, 2009.
- 15 Lanz, V. A., Alfarra, M. R., Baltensperger, U., Buchmann, B., Hueglin, C., and Prévôt, A. S. H.: Source apportionment of submicron organic aerosols at an urban site by factor analytical modelling of aerosol mass spectra, *Atmos. Chem. Phys.*, 7, 1503–1522, doi:10.5194/acp-7-1503-2007, 2007.
- Larsen, B. R., Gilardoni, S., Stenström, K., Niedzialek, J., Jimenez, J., and Belis, C. A.: Sources for PM air pollution in the Po Plain, Italy: II. Probabilistic uncertainty characterization and sensitivity analysis of secondary and primary sources, *Atmos. Environ.*, 50, 203–213, doi:10.1016/j.atmosenv.2011.12.038, 2012.
- 20
- Lee, A. K. Y., Herckes, P., Leaitch, W. R., Macdonald, A. M., and Abbatt, J. P. D.: Aqueous OH oxidation of ambient organic aerosol and cloud water organics: Formation of highly oxidized products, *Geophys. Res. Lett.*, 38, L11805, <https://doi.org/10.1029/2011GL047439>, 2011.
- 25
- Lee, A. K. Y., Hayden, K. L., Herckes, P., Leaitch, W. R., Liggio, J., Macdonald, A. M. and Abbatt, J. P. D.: Characterization of aerosol and cloud water at a mountain site during WACS 2010: secondary organic aerosol formation through oxidative cloud processing, *Atmos. Chem. Phys.*, 12, 7103–7116, <https://doi.org/10.5194/acp-12-7103-2012>, 2012
- 30
- Lenschow, P., Abraham, H. J., Kutzner, K., Lutz, M., Preuss, J. D., and Reichenbacher, W.: Some ideas about the sources of PM10, *Atmos. Environ.*, 35, S23–S33, 2001.

- Li, Y. J., Huang, D. D., Cheung, H. Y., Lee, A. K. Y., and Chan, C. K.: Aqueous-phase photochemical oxidation and direct photolysis of vanillin – a model compound of methoxy phenols from biomass burning, *Atmos. Chem. Phys.*, 14, 2871–2885, <https://doi.org/10.5194/acp-14-2871-2014>, 2014.
- 5 Lim, Y. B., Tan, Y., Perri, M. J., Seitzinger, S. P., and Turpin, B. J.: Aqueous chemistry and its role in secondary organic aerosol (SOA) formation, *Atmos. Chem. Phys.*, 10, 10521–10539, <https://doi.org/10.5194/acp-10-10521-2010>, 2010.
- Matsui, H., Koike, M., Takegawa, N., et al.: Secondary Organic Aerosol Formation in Urban Air: Temporal Variations and Possible Contributions from Unidentified Hydrocarbons, *J. Geophys. Res.*, 114, D04201, doi:10.1029/2008JD010164, 10 2009.
- Matta, E., Facchini, M. C., Decesari, S., Mircea, M., Cavalli, F., Fuzzi, S., Putaud, J.-P., and Dell’Acqua, A.: Mass closure on the chemical species in size-segregated atmospheric aerosol collected in an urban area of the Po Valley, Italy, *Atmos. Chem. Phys.*, 3, 623–637, doi:10.5194/acp-3-623-2003, 2003.
- 15 McNeill, V. F.: Aqueous Organic Chemistry in the Atmosphere: Sources and Chemical Processing of Organic Aerosols, *Environ. Sci. Technol.*, 49, 1237–1244, <https://doi.org/10.1021/es5043707>, 2015.
- Middlebrook, A. M., Bahreini, R., Jimenez, J. L., and Canagaratna, M. R.: Evaluation of composition-dependent collection efficiencies for the Aerodyne aerosol mass spectrometer using field data, *Aerosol Sci. Technol.*, 46, 258–271, doi:10.1080/02786826.2011.620041, 2012.
- 20 Mohr, C., DeCarlo, P. F., Heringa, M. F., Chirico, R., Slowik, J. G., Richter, R., Reche, C., Alastuey, A., Querol, X., Seco, R., Peñuelas, J., Jiménez, J. L., Crippa, M., Zimmermann, R., Baltensperger, U., and Prévôt, A. S. H.: Identification and 25 quantification of organic aerosol from cooking and other sources in Barcelona using aerosol mass spectrometer data, *Atmos. Chem. Phys.*, 12, 1649–1665, doi:10.5194/acp-12-1649-2012, 2012.
- Montero-Martínez, G., Rinaldi, M., Gilardoni, S., Giulianelli, L., Paglione, M., Decesari, S., Fuzzi, S., Facchini, M.C.: On the water-soluble organic nitrogen concentration and mass size distribution during the fog season in the Po Valley, Italy. *Sci. Total Environ.* 485e486, 103e109, 2014.
- 30 Ng, N. L., Canagaratna, M. R., Zhang, Q., Jimenez, J. L., Tian, J., Ulbrich, I. M., Kroll, J. H., Docherty, K. S., Chhabra, P. S., Bahreini, R., Murphy, S. M., Seinfeld, J. H., Hildebrandt, L., Donahue, N. M., DeCarlo, P. F., Lanz, V. A., Prévôt, A. S.

- H., Dinar, E., Rudich, Y., and Worsnop, D. R.: Organic aerosol components observed in Northern Hemispheric datasets from Aerosol Mass Spectrometry, *Atmos. Chem. Phys.*, 10, 4625-4641, <https://doi.org/10.5194/acp-10-4625-2010>, 2010.
- 5 Ng, N. L., Canagaratna, M. R., Jimenez, J. L., Zhang, Q., Ulbrich, I. M., and Worsnop, D. R.: Real-time methods for estimating organic component mass concentrations from aerosol mass spectrometer data, *Environ. Sci. Technol.*, 45, 910-916, 2011.
- Niu, F., Li Z., Li, C., Lee, K-H., and Wang M.: Increase of wintertime fog in China: Potential impacts of weakening of the Eastern Asian monsoon circulation and increasing aerosol loading *J. Geophys. Res.*, 115, doi:10.1029/2009JD013484, 2010.
- 10 Paatero, P. and Tapper, U.: Positive matrix factorization – a nonnegative factor model with optimal utilization of error-estimates of data values, *Environmetrics*, 5, 111–126, doi:10.1002/env.3170050203, 1994.
- Paatero, P., The multilinear engine - A table-driven, least squares program for solving multilinear problems, including the n-way parallel factor analysis model, *Journal of Computational and Graphical Statistics*, 8, 854-888, doi:10.2307/1390831, 1999.
- 15 Paatero, P.: User's guide for the multilinear engine program "ME2" for fitting multilinear and quasimultilinear models, University of Helsinki, Finland, 2000.
- 20 Paglione, M., Saarikoski, S., Carbone, S., Hillamo, R., Facchini, M. C., Finessi, E., Giulianelli, L., Carbone, C., Fuzzi, S., Moretti, F., Tagliavini, E., Swietlicki, E., Eriksson Stenström, K., Prévôt, A. S. H., Massoli, P., Canaragatna, M., Worsnop, D., and Decesari, S.: Primary and secondary biomass burning aerosols determined by proton nuclear magnetic resonance (1H-NMR) spectroscopy during the 2008 EUCAARI campaign in the Po Valley (Italy), *Atmos. Chem. Phys.*, 14, 5089-5110, doi:10.5194/acp-14-5089-2014, 2014.
- 25 Perrone, M. G., Larsen, B. R., Ferrero, L., Sangiorgi, G., De Gennaro, G., Udisti, R., Zangrando, R., Gambaro, A., and Bolzacchini, E.: Sources of high PM_{2.5} concentrations in Milan, northern Italy: molecular marker data and CMB modelling, *Sci. Total Environ.*, 414, 343–355, doi:10.1016/j.scitotenv.2011.11.026, 2012.
- 30 Pietrogrande, M.C., Bacco, D., Visentin, M., Ferrari, S., Poluzzi, V.: Polar organic marker compounds in atmospheric aerosol in the Po Valley during the Supersito campaigns - Part 1: low molecular weight carboxylic acids in cold seasons. *Atmos. Environ.* 86, 164e175, 2014.

- Putaud, J. P., Van Dingenen, R., and Raes, F.: Submicron aerosol mass balance at urban and semirural sites in the Milan area (Italy), *J. Geophys. Res.-Atmos.*, 107(D22), LOP 11–1–LOP 11–10, doi:10.1029/2000JD000111, 2002.
- Putaud, J. P., Van Dingenen, R., Alastuey, A., Bauer, H., Birmili, W., Cyrys, J., Flentje, H., Fuzzi, S., Gehrig, R., Hansson, H. C., Harrison, R. M., Herrmann, H., Hittenberger, R., Hueglin, C., Jones, A. M., Kasper-Giebl, A., Kiss, G., Kousa, A., Kuhlbusch, T. A. J., Loeschau, G., Maenhaut, W., Molnar, A., Moreno, T., Pekkanen, J., Perrino, C., Pitz, M., Puxbaum, H., Querol, X., Rodriguez, S., Salma, I., Schwarz, J., Smolik, J., Schneider, J., Spindler, G., ten Brink, H., Tursic, J., Viana, M., Wiedensohler, A., and Raes, F.: A European aerosol phenomenology-3: Physical and chemical characteristics of particulate matter from 60 rural, urban, and kerbside sites across Europe, *Atmos. Environ.*, 44, 1308–1320, 2010.
- Ricciardelli, I., Bacco, D., Rinaldi, M., Bonafe, G., Scotto, F., Trentini, A., Bertacci, G., Ugolini, P., Zigola, C., Rovere, F., Maccone, C., Pironi, C., and Poluzzi, V.: A three-year investigation of daily PM 2.5 main chemical components in four sites: the routine measurement program of the Super-sito Project (Po Valley, Italy), *Atmos. Environ.*, 152, 418–430, <https://doi.org/10.1016/j.atmosenv.2016.12.052>, 2017.
- Rinaldi, M., Gilardoni, S., Paglione, M., Sandrini, S., Fuzzi, S., Massoli, P., Bonasoni, P., Cristofanelli, P., Marinoni, A., Poluzzi, V., and Decesari, S.: Organic aerosol evolution and transport observed at Mt. Cimone (2165 m a.s.l.), Italy, during the PEGASOS campaign, *Atmos. Chem. Phys.*, 15, 11327–11340, <https://doi.org/10.5194/acp-15-11327-2015>, 2015.
- Saarikoski, S., Carbone, S., Decesari, S., Giulianelli, L., Angelini, F., Canagaratna, M., Ng, N. L., Trimborn, A., Facchini, M. C., Fuzzi, S., Hillamo, R., and Worsnop, D.: Chemical characterization of springtime submicrometer aerosol in Po Valley, Italy, *Atmos. Chem. Phys.*, 12, 8401–8421, doi:10.5194/acp-12-8401-2012, 2012.
- Sandrini, S., van Pinxteren, D., Giulianelli, L., Herrmann, H., Poulain, L., Facchini, M. C., Gilardoni, S., Rinaldi, M., Paglione, M., Turpin, B. J., Pollini, F., Bucci, S., Zanca, N., and Decesari, S.: Size-resolved aerosol composition at an urban and a rural site in the Po Valley in summertime: implications for secondary aerosol formation, *Atmos. Chem. Phys.*, 16, 10879–10897, <https://doi.org/10.5194/acp-16-10879-2016>, 2016.
- Saraf, A. K., Bora, A. K., Das, J., Rawat, V., Sharma, K., Jain, S. K.: Winter fog over the Indo-Gangetic Plains: Mapping and modelling using remote sensing and GIS, *Nat. Hazards*, 58, 199–220, DOI 10.1007/s11069-010-9660-0, 2011.
- [Schauer J.J., Kleeman M.J., Cass G.R., Simoneit B.R.T.: Measurement of emissions from air pollution sources. 3. C1-C29 organic compounds from fireplace combustion of wood. Environ Sci Technol 35\(9\):1716–1728, 2001.](#)

- Schneider, J., Weimer, S., Drewnick, F., Borrmann, S., Helas, G., Gwaze, P., Schmid, O., Andreae, M. O. and Kirchner, U.: Mass spectrometric analysis and aerodynamic properties of various types of combustion-related aerosol particles, *Int. J. Mass. Spec.*, 258, 37–49, 2006.
- 5 Stanier, C., Singh, A., Adamski, W., Baek, J., Caughey, M., Carmichael, G., Edgerton, E., Kenski, D., Koerber, M., Oleson, J., Rohlf, T., Lee, S. R., Riemer, N., Shaw, S., Sousan, S., and Spak, S. N.: Overview of the LADCO winter nitrate study: hourly ammonia, nitric acid and PM_{2.5} composition at an urban and rural site pair during PM_{2.5} episodes in the US Great Lakes region, *Atmos. Chem. Phys.*, 12, 11037-11056, <https://doi.org/10.5194/acp-12-11037-2012>, 2012.
- 10 Stavroulas, I., Bougiatioti, A., Paraskevopoulou, D., Grivas, G., Liakakou, E., Gerasopoulos, E., and Mihalopoulos, N.: Sources and processes that control the submicron organic aerosol in an urban Mediterranean environment (Athens) using high temporal resolution chemical composition measurements, *Atmos. Chem. Phys. Discuss.*, <https://doi.org/10.5194/acp-2018-356>, in review, 2018.
- 15 Sullivan, A. P., Hodas, N., Turpin, B. J., Skog, K., Keutsch, F. N., Gilardoni, S., Paglione, M., Rinaldi, M., Decesari, S., Facchini, M. C., Poulain, L., Herrmann, H., Wiedensohler, A., Nemitz, E., Twigg, M. M., and Collett Jr., J. L.: Evidence for ambient dark aqueous SOA formation in the Po Valley, Italy, *Atmos. Chem. Phys.*, 16, 8095-8108, <https://doi.org/10.5194/acp-16-8095-2016>, 2016.
- 20 Sun, Y. L., Zhang, Q., Anastasio, C., and Sun, J.: Insights into secondary organic aerosol formed via aqueous-phase reactions of phenolic compounds based on high resolution mass spectrometry, *Atmos. Chem. Phys.*, 10, 4809-4822, <https://doi.org/10.5194/acp-10-4809-2010>, 2010.
- Sun, Y., Xu, W., Zhang, Q., Jiang, Q., Canonaco, F., Prévôt, A. S. H., Fu, P., Li, J., Jayne, J., Worsnop, D. R., and Wang, Z.: Source apportionment of organic aerosol from 2-year highly time-resolved measurements by an aerosol chemical speciation monitor in Beijing, China, *Atmos. Chem. Phys.*, 18, 8469-8489, <https://doi.org/10.5194/acp-18-8469-2018>, 2018.
- 30 [Timonen, H., Carbone, S., Aurela, M., Saarnio, K., Saarikoski, S., Ng, N.L., Canagaratna, M.R., Kulmala, M., Kerminen, V.-M., Worsnop, D.R., Hillamo, R.: Characteristics, sources and water-solubility of ambient submicron organic aerosol in springtime in Helsinki, Finland, *J. Aerosol Sci.*, 56, pp. 61-77, \[10.1016/j.jaerosci.2012.06.005\]\(https://doi.org/10.1016/j.jaerosci.2012.06.005\), 2013.](#)
- Tsimpidi, A. P., Karydis, V. A., Pandis, S. N., and Lelieveld, J.: Global combustion sources of organic aerosols: model comparison with 84 AMS factor-analysis data sets, *Atmos. Chem. Phys.*, 16, 8939-8962, <https://doi.org/10.5194/acp-16-8939-2016>, 2016.

- Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.: Interpretation of organic components from Positive Matrix Factorization of aerosol mass spectrometric data, *Atmos. Chem. Phys.*, 9, 2891–2918, doi:10.5194/acp-9-2891-2009, 2009.
- 5
- Ulbrich, I. M., Lechner, M., and Jimenez, J. L.: AMS Spectral Database, available at: <http://cires.colorado.edu/jimenez-group/AMSsd/>, last access: 7 October 2015.
- Whiteaker, J. R., and Prather K. A.: Hydroxymethanesulfonate as a tracer for fog processing of individual aerosol particles. *10 Atmospheric Environment*. 37:1033-1043, 2003.
- WHO, European Detailed Mortality Database, Update July 2016. WHO Regional Office for Europe, Copenhagen. <http://data.euro.who.int/dmdb/>, 2016.
- 15 Wu, X., Vu, T.V., Shi, Z., Harrison, R.M., Liu, D., and Cen, K.: Characterization and source apportionment of carbonaceous PM_{2.5} particles in China - a review, *Atmos. Environ.*, 189, pp. 187-212, 2018
- Xu, L., Suresh, S., Guo, H., Weber, R. J., and Ng, N. L.: Aerosol characterization over the southeastern United States using high-resolution aerosol mass spectrometry: spatial and seasonal variation of aerosol composition and sources with a focus on organic nitrates, *Atmos. Chem. Phys.*, 15, 7307-7336, <https://doi.org/10.5194/acp-15-7307-2015>, 2015.
- 20
- Xu, W., Han, T., Du, W., Wang, Q., Chen, C., Zhao, J., Zhang, Y., Li, J., Fu, P., Wang, Z., Worsnop, D. R., and Sun, Y.: Effects of Aqueous-Phase and Photochemical Processing on Secondary Organic Aerosol Formation and Evolution in Beijing, China, *Environ. Sci. Technol.*, 51 (2), 762-770, DOI: 10.1021/acs.est.6b04498, 2017.
- 25
- Yu, L., Smith, J., Laskin, A., Anastasio, C., Laskin, J., and Zhang, Q.: Chemical characterization of SOA formed from aqueous-phase reactions of phenols with the triplet excited state of carbonyl and hydroxyl radical, *Atmos. Chem. Phys.*, 14, 13801-13816, <https://doi.org/10.5194/acp-14-13801-2014>, 2014.
- 30 Zhang Q., Jimenez J. L., Canagaratna M. R., Allan J. D., Coe H., Ulbrich I., Alfarra M. R., Takami A., Middlebrook A. M., Sun Y. L., Dzepina K., Dunlea E., Docherty K., DeCarlo P. F., Salcedo D., Onasch T., Jayne J. T., Miyoshi T., Shimo A., Hatakeyama S., Takegawa N., Kondo Y., Schneider J., Drewnick F., Borrmann S., Weimer S., Demerjian K., Williams P., Bower K., Bahreini R., Cottrell L., Griffin R. J., Rautiainen J., Sun J. Y., Zhang Y. M., and Worsnop D. R.: Ubiquity and

dominance of oxygenated species in organic aerosols in anthropogenically-influenced Northern Hemisphere midlatitudes. *Geophys Res Lett* 34(13):L13801, 2007.

5 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Ulbrich, I. M., Ng, N. L., Worsnop, D. R., and Sun, Y.: Understanding atmospheric organic aerosols via factor analysis of aerosol mass spectrometry: a review, *Anal. Bioanal. Chem.*, 401, 3045–3067, doi:10.1007/s00216-011-5355-y, 2011.

Table 1: Average organic aerosol (OA) concentrations and its relative contribution to the NR-PM1 mass measured by the HR-TOF-AMS within each campaign.

	BO		SPC	
	OA ($\mu\text{g m}^{-3}$)	OA/NR-PM1	OA ($\mu\text{g m}^{-3}$)	OA/NR-PM1
2011 Fall	15.85	46%	9.30	50%
2012 Summer	7.16	58%	5.27	49%?
2012 Fall	4.61	46%		
2013 Winter	8.37	42%		
2013 Spring	2.04	44%	1.74	36%
2013 Fall	3.81	33%	3.37	40%
2014 Winter	3.60	39%		
2014 Spring	3.31	54%		

5 **Table 2: Relative (%) and absolute mass contribution ($\mu\text{g m}^{-3}$) of main organic aerosol components HOA, BBOA, COA and OOA for all the considered campaigns. BO = Bologna, SPC = San Pietro Capofiume.**

			HOA	BBOA	COA	OOA	
BO	SPRING	2013	12% (0.25)	14% (0.29)	-	73% (1.49)	
		2014	6% (0.18)	2% (0.06)	8% (0.28)	84% (2.71)	
	SUMMER	2012	8% (0.58)	-	-	92% (6.58)	
		FALL	2011	18% (2.80)	38% (6.05)	-	44% (7.00)
			2012	16% (0.74)	30% (1.37)	-	54% (2.50)
	WINTER	2013	11% (0.43)	17% (0.64)	-	72% (2.74)	
		2014	11% (0.88)	28% (2.35)	-	61% (5.14)	
	SPC	SPRING	2013	9% (0.15)	3% (0.05)	-	88% (1.54)
2012			4% (0.20)	-	-	96% (5.06)	
FALL		2011	32% (2.93)	33% (3.07)	-	35% (3.29)	
		2013	7% (0.23)	28% (0.95)	-	65% (2.20)	

Table 3: Urban increment, calculated as the ratio between the campaign average concentration in urban and rural site, for each season and OA fraction considered.

Urban Increment	HOA	BBOA	OOA	OA TOT
SPRING 2013	1.67	5.87	0.97	1.17
SUMMER 2012	2.85	-	1.30	1.36
FALL 2013	1.91	0.67	1.25	1.13

5

Table 4: Correlation (Pearson coefficients, R) between OOA components identified by PMF/ME-2 and some variables linked to aqueous phase: Relative Humidity (RH) of the air; Aerosol Liquid Water Content (ALWC) calculated by the ISORROPIA II model; hydroxymethanesulfonate (HMSA) estimated by AMS and (when available) NMR measurements. The shaded cells highlight the highest correlations with a color scale ranging from less to more intense orange as the R value increases. The light-blue shaded cells highlight the identified aqSOA factor. Gray cells indicate missing data.

10

			BO					SPC			
			RH	ALWC	HMSA (AMS)	HMSA (NMR)		RH	ALWC	HMSA (AMS)	HMSA (NMR)
SPRING	2013_spring (may)	OOA1_BB	0.53	0.56	0.12	-	OOA1	0.38	0.65	0.36	
		OOA2	-0.15	0.53	0.40	-	OOA2	0.19	0.76	0.55	
		OOA3_BB-aq	0.33	0.75	0.62	-	OOA3	-0.09	0.53	0.58	
	2014_spring (may)	OOA1	0.25	0.11	-0.03	-					
		OOA2	-0.09	-0.07	0.40	-					
		OOA3	0.13	0.23	0.56	-					
SUMMER	2012_summer (jun.-jul.)	OOA1	-0.26	0.12	-0.28	0.32	OOA1	0.08	0.20		0.33
		OOA2	0.38	0.31	-0.47	0.13	OOA2	0.50	0.79		0.20
							OOA3	0.15	0.38		0.51
							OOA4	-0.27	-0.22		0.42
FALL	2011_fall. (nov.-dic.)	OOA1_BB	-0.12	0.63	-0.06	-	OOA_BB-aq	0.00	0.88	0.77	0.66
		OOA2_BB-aq	0.43	0.82	0.58	-					
	2012_fall (oct.-nov.)	OOA1	0.00	-0.01	-0.04	-					
		OOA2_BB-aq	0.26	0.83	0.70	-					
		OOA3_BB	0.11	0.20	0.55	-					
	2013_fall (oct.)	OOA1	-0.29	0.16	-0.22	0.15	OOA1_BB	-0.24	-0.29	-0.06	-0.17
		OOA2	0.28	0.65	0.76	0.53	OOA2_BB-aq	0.07	0.63	0.68	0.70
		OOA3_BB-aq	0.34	0.82	0.81	0.86	OOA3	-0.04	0.28	0.45	0.25
	WINTER	2013_winter (jan.-feb.)	OOA1_BB	0.06	0.43	0.24	0.23				
OOA2_BB-aq			0.33	0.73	0.68	0.47					
OOA3			0.29	0.32	0.19	0.16					
2014_winter (jan.-feb.)		OOA1_BB	0.16	0.38	0.43	0.57					
		OOA2_BB-aq	0.31	0.74	0.72	0.71					
		OOA3	-0.11	0.63	0.44	0.60					

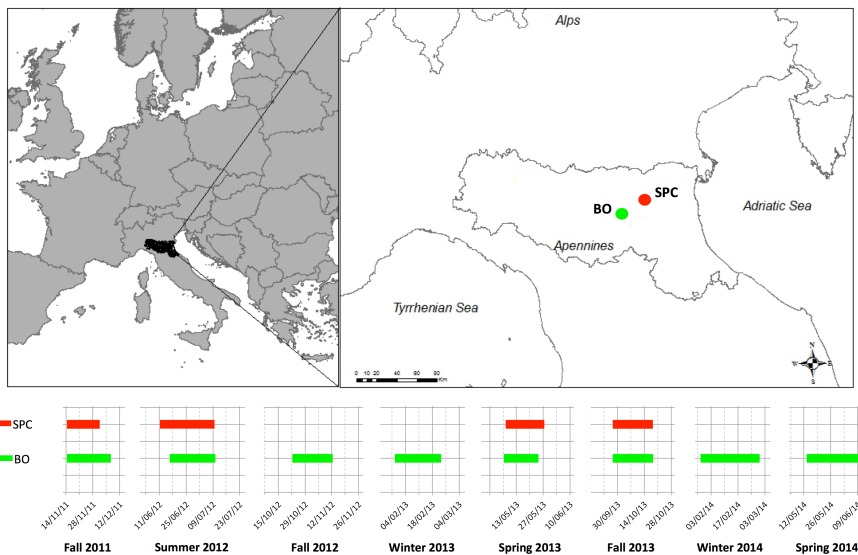
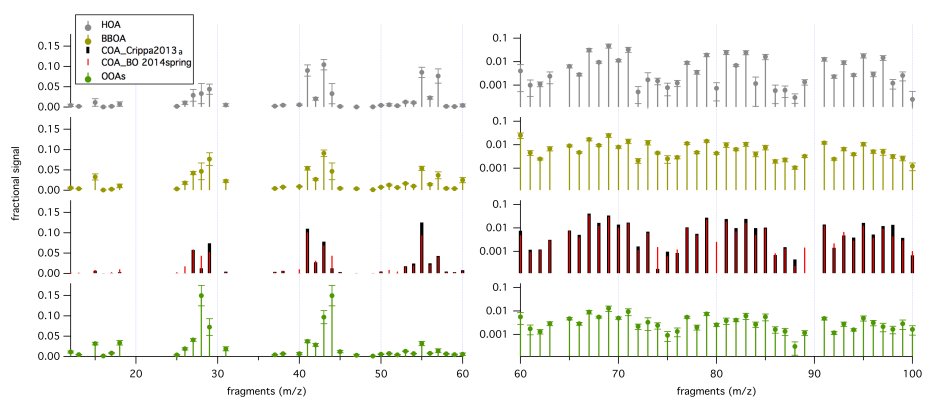


Figure 1: SUPERSITO field campaigns: map of the sites and measurement periods considered in this study.



5 Figure 2: Mass spectral variability for the main retrieved OA sources. Mean values are represented with circles and the \pm standard deviation with error bars. COA from the BO spring 2014 campaign is represented in red color over-imposed to the COA reference spectrum from Crippa et al. 2013a.

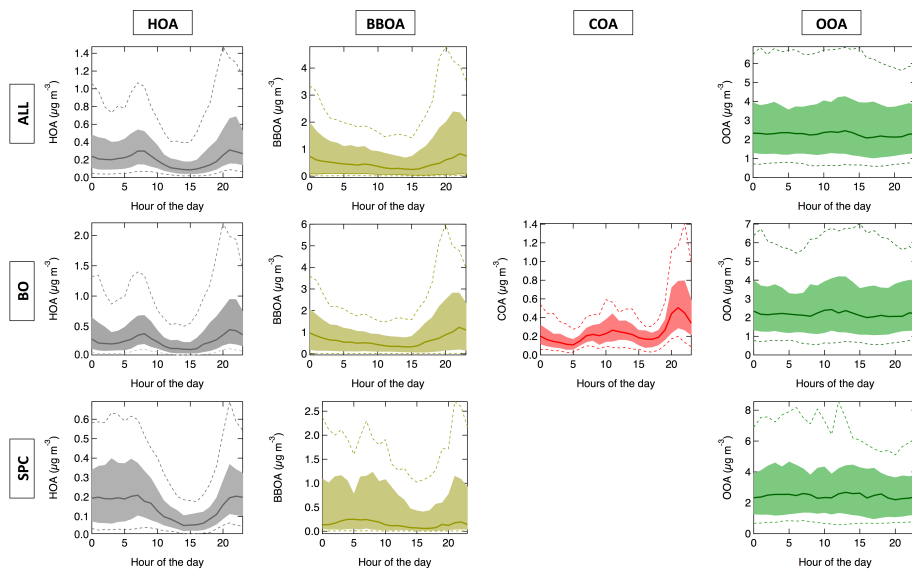


Figure 3: Daily trends of the factors identified (HOA, BBOA, COA and SOA). Median diurnal pattern are reported together with the 10th, 25th, 75th and 90th percentiles for each source, for all the Supersito campaigns together and separately for Bologna (BO) and San Pietro Capofiume (SPC).

5

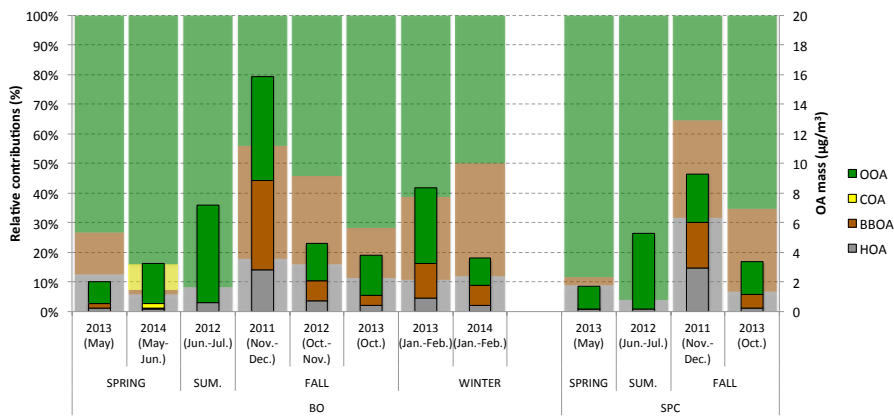


Figure 4: Organic aerosol sources contribution for each site and each SUPERSITO campaign. Relative contributions are reported as shaded histograms (referring to the left axis) in the background of the absolute ones (referring to right axis).

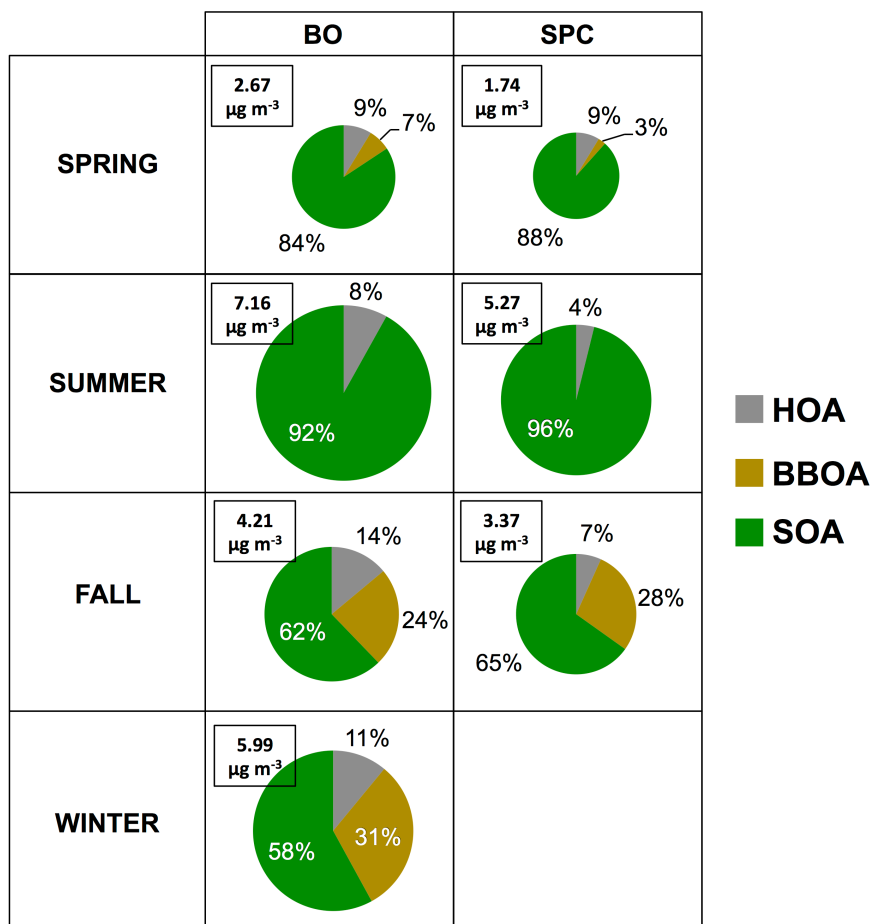


Figure 5: Seasonal relative contribution of the main OA sources in both urban and rural site. Pie-charts area is proportional to the total average concentration of OA (reported in the upper-left side of each box in term of $\mu\text{g m}^{-3}$) and the individual portions are the average between the different campaigns made in the site in one season.

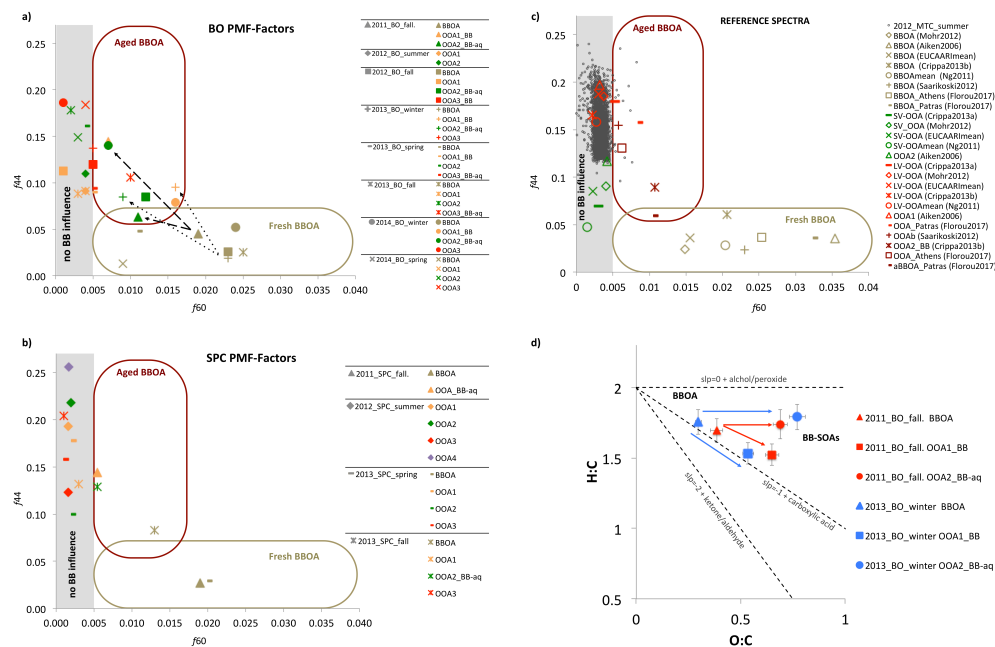


Figure 6: Influence of Biomass Burning emissions on SOA and their evolution processes. The plots in panel a, b and c show f_{44} (normalized mass spectrum at m/z 44), which is a proxy of OA oxygenation degree, versus f_{60} (normalized mass spectrum signal at m/z 60), which is a proxy of anhydrosugars. Different shapes of the markers identify different SUPERSITO campaigns (panel a and b) or different reference spectra (panel c). Different colors represents the different kind of PMF-factors: gold-green identifies BBOA primary factors, yellow, green and red the OOA numerically ordered based on their O:C ratios. Black dots in panel c represent the measurements taken as background level of no influence of biomass burning. Gray areas correspond to $f_{60} 0.003 \pm 0.002$ representing the Cubison et al. 2010 threshold of BB influence. Panel d) reports Van Krevelen (VK) diagram of the BBOA and OOA-BB PMF factors obtained from the HR-ToF-AMS data analysis for both BO fall 2011 (red markers) and winter 2013 (blue markers). The line connecting BBOA and OOA-BB has different slopes, indicating different chemistry processing leading to different SOA types.

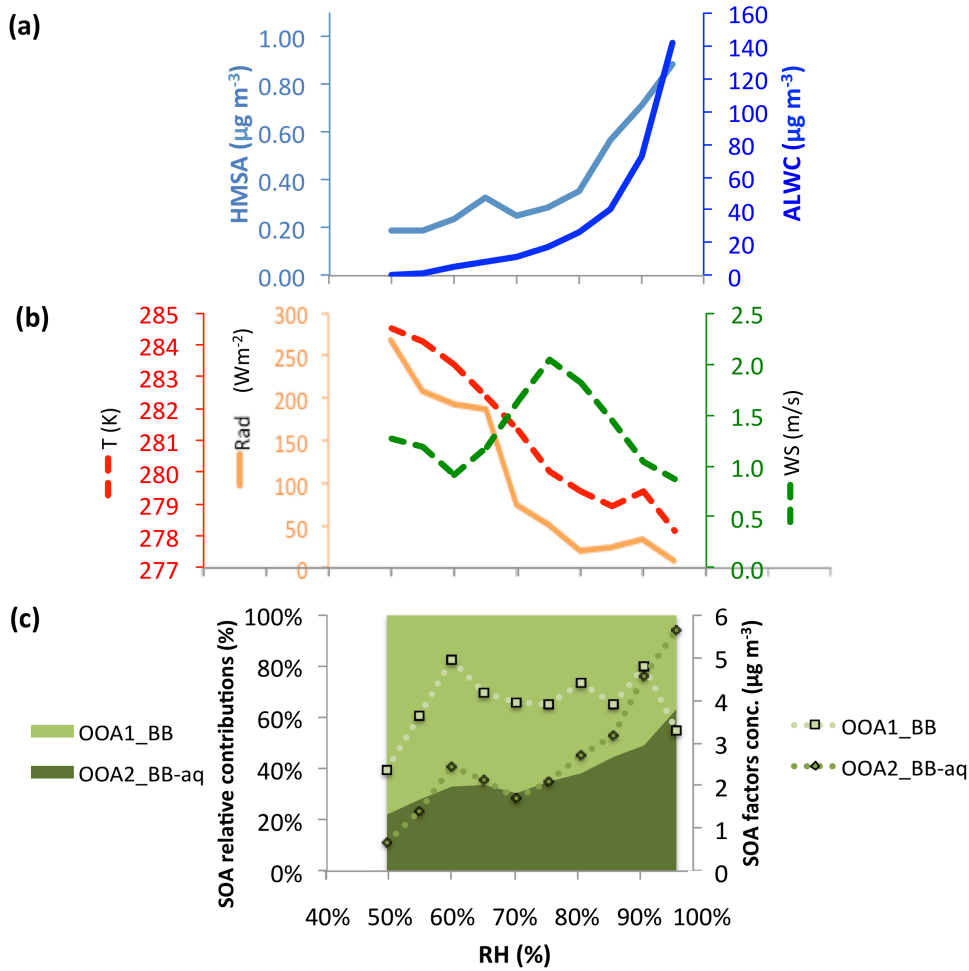


Figure 7: variations of meteorological and chemical parameters as function of RH during the BO fall 2011 campaign. The data were binned according to the RH (5% increment), and mean values are shown for each bin. Panel (a): Aerosol Liquid Water Content (ALWC) and hydroxymethansulfonic acid (HMSA). Panel (b): air temperature together with solar radiation and wind speed (WS) measured at ground level. Panel (c): variations in contributions of the two BB-influenced OOA factors identified (OOA1_BB and OOA2_BB-aq) both in absolute ($\mu\text{g m}^{-3}$) and relative (% of OOA) terms.

- Marco Paglione 5/8/y 17:16
Eliminato: 10
- Marco Paglione 5/8/y 16:33
Eliminato: A
- Marco Paglione 5/8/y 16:33
Eliminato: B
- Marco Paglione 5/8/y 16:33
Eliminato: C
- Marco Paglione 5/8/y 16:33
Eliminato: Panel D: different SOA categories excluding the effects of planetary boundary layer height (PBL) using the total POA as a surrogate.

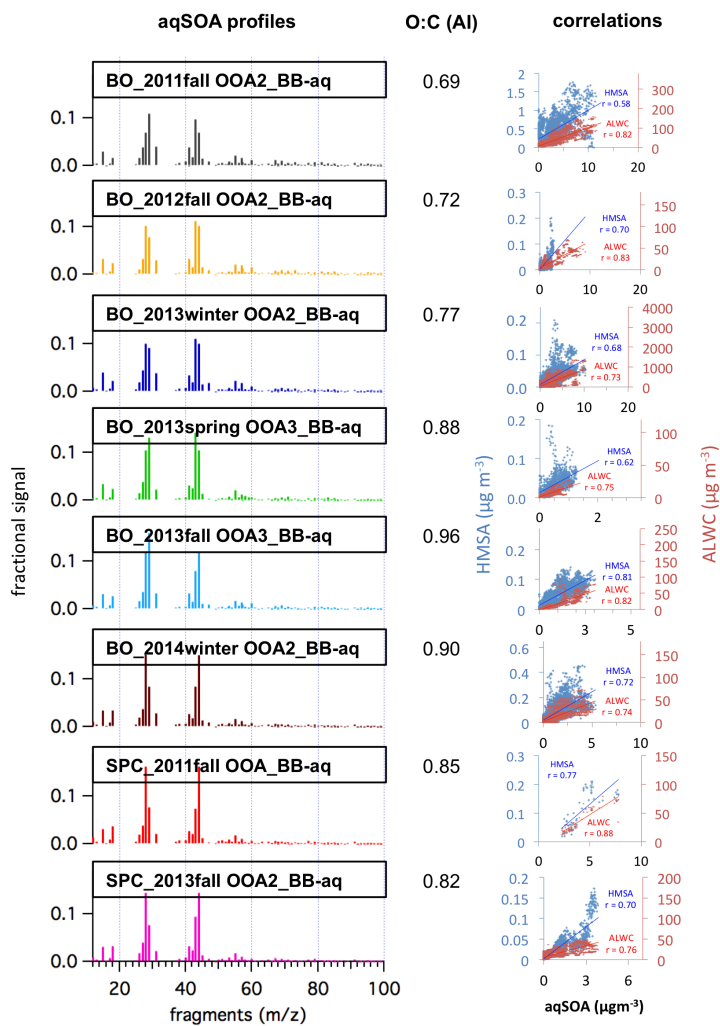


Figure 8: OOA_x BB-aq main features: left column shows the mass spectral profile of each BB-aqSOA component identified during the SUPERSITO campaigns; the OOA factors are numerically ordered for each campaign based on their O:C ratios; central column reports the O:C elemental ratios of the same factors; right column illustrates the correlation between their concentration time series and the HMSA (in blue) and the ALWC (in red).

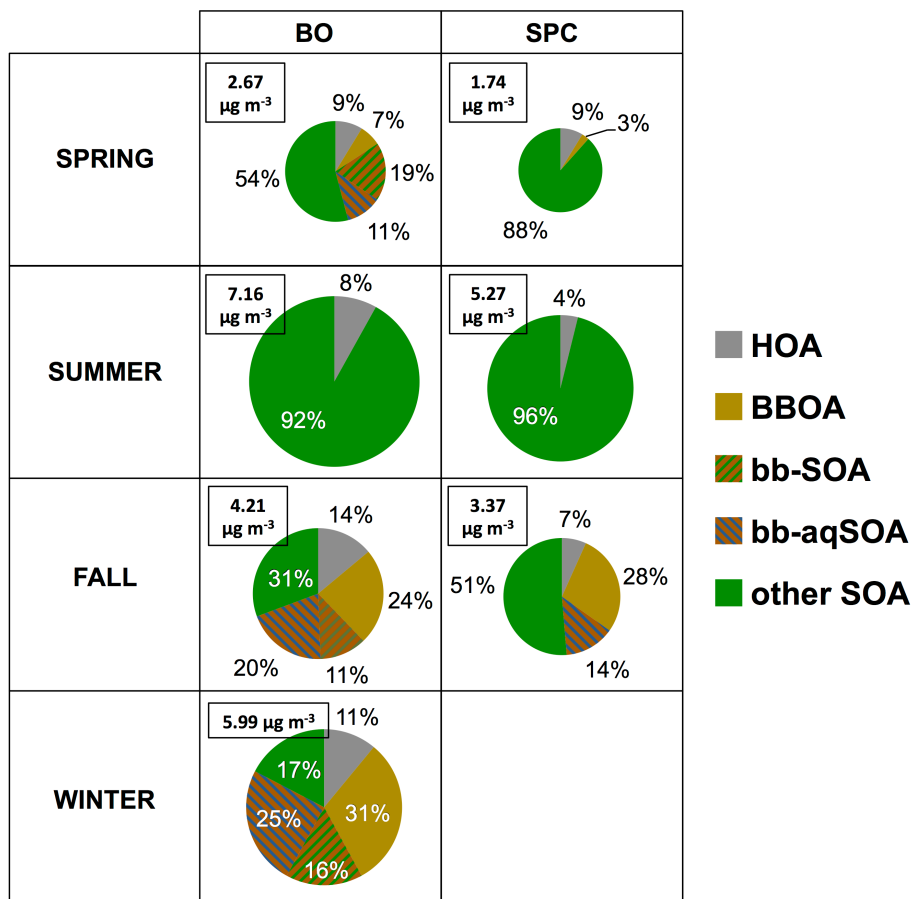


Figure 9: Seasonal relative contribution of the main OA sources in both urban and rural site with explicit separation also of the SOA (OOA) components. Pie-charts area is proportional to the total average concentration of OA (reported in the upper-left side of each box in term of $\mu\text{g m}^{-3}$) and the individual portions are the average between the different campaigns made in the site in one season. OOA factors influenced by biomass burning (characterized by brown background color) are divided in the two categories, "bb-SOA" and "bb-aqSOA" representing the OOA_{x_BB} and OOA_{x_BB-aq} described in the text. "Other SOA" is the sum of the other OOA factors whose source has not been unequivocally identified.

Marco Paglione 5/8/y 17:16

Formattato: Didascalia