Replies to Referee Comments (RC1, RC2) and Short Comment (SC1)

² We thank both referees No.1 and No.2 for their detailed and insightful comments which helped to ³ further improve the manuscript. Below we address all referee comments as well as the contributed ⁴ short comment (in italics). Revised text, keyed to the ACPD online version, is shown in blue, and ⁵ is included in the final manuscript we are submitting to ACP. Updated figures and tables as well as ⁶ material for a supplement are attached at the end of our reply.

7

RC1: Source assignment of proxies is a basic prerequisite for interpreting climate archives in terms 8 of past climate as well as climate change. Concerning polar ice cores, ionic impurities originate pri-9 marily from aerosol deposition. Amongst them, interpretation of sea salt aerosol deposition archived 10 in ice cores is especially challenging and contro- versial because the contribution of two different and 11 competing sources - viz. open water versus sea ice - is up for debate. In addition it became apparent 12 that sea salt aerosol production over ice-covered oceans may contribute significantly to the global sea 13 salt aerosol budget. The manuscript at hand addresses this pivotal subject and provides thorough 14 and direct observational evidence of sea salt aerosol production from blowing snow above sea ice. 15 The important conclusions drawn are based on comprehensive state of the art ship-borne aerosol 16 and snow measurements during winter / early springtime in the Weddell Sea region. Although the 17 main conclusions are primarily restricted to the chosen site, there are certainly strong implications for 18 climate research in the Southern Ocean realm and climate related interpretation of sea salt profiles 19 from ice cores in general. The authors have accomplished a clear, well-organised and concise paper. 20 The methodology is sound and assumptions are identified clearly and conscientiously. From my point 21 of view, all parts, including figures, are essential. The manuscript certainly addresses the scientific 22

23 scope of ACP and I recommend a final publication after some minor revisions I specified below.

24 **Reply:** We thank the reviewer for the positive assessment of this work.

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RC1: 1. Abstract, page 1, line 13 and Conclusions, page 24, line19: The authors state that bromine enrichment was typical at 29 m height, but from Chapter 3.4.2 and Fig. 15, bromine depletion is evident. Please clarify.

²⁹ **Reply:** In both cases it should read "depleted", which is now corrected.

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- **RC1:** 2. Chapter 2.3, Aerosol chemical composition: Could you assess the impact of pollution on chemical aerosol composition? Was the bulk aerosol sampling contamination controlled?
- **Reply:** The setup of this study did not include any contamination control of the bulk aerosol filter sampling such as pump control based on wind speed and direction. As described aerosol number concentrations at the crows nest showed significant spikes, when air came from the direction of the ship stack, whereas no evidence of pollution was detected in aerosol number concentrations observed on the sea ice. We clarified text in section 2.2 including also the fraction of CLASP data
- filtered out:
 Raw aerosol number concentrations at the crow's nest showed significant spikes, when air came from
- the direction of the ship's engine stack, whereas no evidence of pollution was detected in the obser-
- vations on the sea ice. Pollution spikes were effectively filtered out prior to averaging by excluding all
 data when relative wind direction was in the 135–225° sector encompassing the ship's engine stack.
- A total of 21% of the available 1-second data was removed from the crow's nest data.
- We now include an assessment of the potential impact of pollution on bulk aerosol chemistry from filters and added the following text to section 2.3:
- ⁴⁶ In order to assess the impact of potential pollution on bulk aerosol chemistry from filters collected at
- the crow's nest we calculated for each filter sample the fraction of the total filter run time during which
- relative wind direction was within the 135–225° sector encompassing the ship's engine stack. Con-
- 49 sidering all filters sampled from June to September 2013 (N=141) the fraction of total filter run time
- ⁵⁰ with winds from the polluted sector was on average 9.5%. Polluted time fraction and atmospheric
- ⁵¹ concentrations of Na⁺, Cl⁻, SO₄²⁻ and Br⁻ did not show any correlation (R²<0.05), suggesting that
- ⁵² the impact of pollution on the respective ion concentrations is small. A weak, but significant negative
- correlation was found between polluted time fraction and depletion factors $DF_{SO_4^{2-}}$ (R²=0.19, p<0.01)

and DF_{Br^-} (R²=0.13, p<0.01) suggesting that enrichment in sulphate (and bromide) may be more likely during polluted conditions. The bulk aerosol chemistry observations on the sea ice showed no evidence of pollution. Thus, in the case of sulfate we cannot rule out that some of the sulfate enrichment in atmospheric aerosol observed at the crow's nest may be due to ship exhaust rather than presence of mirabilite. It follows that estimates of sea ice contributions to total SSA derived from depletion factors discussed in section 3.4.4 have to be considered as lower bounds of true values.

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RC1: 3. Chapter 3.2: Impact of snow precipitation on blowing/drifting snow: Did you access the regular weather reports from the ships meteorological office in this case?

Reply: The available 3-hourly weather reports from the ship provided only limited information, but confirmed for the case of precipitation shown in Fig.10 (3-4 July 2013) overcast skies, variable visibility from 0.5 (fog) to 10 km, and occasional ice needles. As stated in the text we inferred occurrence of precipitation qualitatively from direct observation supported by webcam images, if usable, and presence of clouds (p9-I14). The still images of a webcam installed at the crows nest (p8-I16) allowed to see at times, including 4 July, large airborne snow crystals during night time in the beam of the ships search lights. We added a sentence in section 3.3.2 to further clarify.

⁷⁰ For the early morning of 4 July 2013 webcam images from the crow's nest confirmed the presence of

⁷¹ large airborne snow crystals visible during darkness in the beam of the ship's search lights, whereas

⁷² the ship's 3-hourly weather report noted the presence of airborne ice needles.

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RC1: 4. Pages 12/13 and Fig.7: Regarding the salinity (S_p) of blowing snow, corresponding S_p -

values of the uppermost surface snow layer are decisive. Did you take samples from surface snow;

⁷⁶ say <1 cm deepness below surface? Figure 7: The reader cannot get an idea about the salinity of

the surface snow layer from this graph. It would be informative as well to specify the total depth of

⁷⁸ the snow layer shown here, not just the snow height above sea ice.

Reply: Typically snow pit profiles were measured at 2 cm depth resolution (see methods section 2.4); except at ice station S6 some profiles include a surface snow sample from a layer of ~0.5-1.0 cm thickness. As discussed, S_p in blowing snow is consistent with the local surface snow measurements (section 3.2.2). In Figure 7 depth information is readily available since data points at the top of each profile, i.e. with the largest snow height above ice, represent the surface snow layer. We updated Figure 7 and caption accordingly (Fig. 1).

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RC1: *5.* Chapter 3.3.2, Snow particle size distribution: Is it possible to rate the impact of the ships profile on the local wind field and eventually on the measured snow particle size distribution?

Reply: We did not attempt to quantify the distortion of the local wind field by the ship and its impact
 on measured snow particles, and had therefore included a respective caveat in the method descrip tion (p4 line 30): "It should be borne in mind that the distortion of flow caused by the ship may mean

both that speed at 39 m is not representative of flow in the far field at that height, and further, the turbulent field strength, which governs the gradient of the logarithmic profile, may be a residual from a

different, likely lower, height. Thus, we suggest care when interpreting the data, and estimate that the
 conversion from particle counts to number density be seen as an estimate suitable for comparison,

⁹⁵ rather than quantitative with a well behaved uncertainty."

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RC1: 6. Chapter 3.4.1, page 19, lines 28-30: As for Antarctic winter, acid induced C[⊢] loss is rather extraordinary because production of acidic sulphur compounds usually cease at the end of summer / fall. Are there any indications for alternative HNO₃ induced C[⊢] loss in your data?

Reply: We agree. Sea salt reaction with atmospheric HNO₃ is a plausible alternative chloride loss process in winter; e.g. at Halley on the nearby Antarctic coast observations in winter show low but non-zero levels of atmospheric HNO₃ of 1-2 pptv (Jones et al., 2011). Unfortunately, no usable filter data of aerosol nitrate are available from this study to test the suggested process due to a very high lab procedure blanc. However, we include the point in section 3.4.1 as follows.

Snow on sea ice follows closely the theoretical mirabilite fractionation line, whereas aerosol shows
 large scatter and a tendency to apparent Na⁺ enrichment with respect to Cl⁻ of up to 20 %, equivalent

to Cl⁻ depletion with respect to Na⁺ of 17% (Fig. 14). Dechlorination of sea salt aerosol observed 107 in Antarctica has a maximum in spring/summer, when gaseous acidic species (nitric, sulfuric and 108 methanesulfonic acid) are available to replace chloride on sea-salt aerosol (Wagenbach et al., 1998; 109 Rankin and Wolff, 2003; Legrand et al., 2017). Acidic sulphur species are close to zero during winter 110 in coastal Antarctica e.g. at Neumayer (Weller et al., 2011), whereas nitric acid is low but non-zero, 111 e.g. 1-2 pptv at Halley (Jones et al., 2011). Thus nitric acid induced Cl⁻ loss from sea salt is a plausi-112 ble explanation for the observed CI⁻ depletion either in airborne SSA or as a sampling artefact from 113 sea salt already accumulated on the filter surface as suggested previously (Wagenbach et al., 1998; 114 Legrand et al., 2017). Unfortunately no usable filter data of aerosol nitrate are available from this 115 study to further test the association between nitrate and sea salt due to a very high lab procedure 116 blank. 117 118

RC1: 7. Chapter 3.4.2, Chemical fractionation of Br⁻, lines 28-33 and Fig. 15: There is strong bromine depletion during polar night in July when global radiation was about zero (Fig. 15b). This peculiarity deserves some discussion.

122 **Reply:** We did mention (p21-line16) that bromide escape from aerosol was detected previously year-

round at DDU in coastal Antarctica, including during winter months, except in June (Legrand et al.,

¹²⁴ 2016). As suggested, we expand the discussion in section 3.4.2:

¹²⁵ Contrary to expectation bromide depletion of aerosol was significant even during winter darkness ¹²⁶ from mid June to mid July (Fig. 15b), whereas previous observations at DDU showed a similar trend ¹²⁷ but less bromide depletion and none in June (Legrand et al., 2016). At DDU DF_{Br^-} in bulk aerosol

increased gradually from a minimum in June (0.04), intermediate values in July to Sep (0.22-0.39) to a maximum in October (0.42) (Legrand et al., 2016). Light conditions are unlikely a cause of differ-

ences in bromide depletion, since DDU is located at a similar latitude (66° 40'S) as the area covered
 by this study. However, one of a number of processes identified leading to bromide loss from snow

¹³² or aerosol involves HOBr oxidation of bromide, which leads to its autocatalytic release (Abbatt et al.,

¹³³ 2012). The early laboratory study by Oum et al. (1998) has shown that the required HOBr can be

chemically produced in darkness through the reaction of ozone with bromide. Another study during
 the ANT-XXIX/6 expedition reports significant bromoform (CH₃Br) production in sea ice during winter

the AN I-XXIX/6 expedition reports significant bromoform (CH₃Br) production in sea ice during winter
 darkness (Abrahamsson et al., 2018), which requires HOBr (and organic matter) as precursors, and
 therefore indicates that bromine loss processes were active in the sea ice in the absence of sunlight.

137 Incretore indicates that bromine loss processes were active in the sea ice in the absence of sunlight.
 138 It therefore appears plausible that the same reactions may have caused significant bromide depletion

observed here in sea salt aerosol, provided the aerosol pH was low enough.

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RC1: 8. Figure 5 and page 12, lines 8-9: By the way: During late afternoon of the 11 July, there is an outstanding Na⁺ peak associated with corresponding sulphate depletion, while the wind speed seemed just close to the threshold value (well below 10 m/s throughout the whole day). Any ideas?

Reply: From midnight to the early morning of 11 July 2013 wind speed was indeed at or slightly 144 above the snow drift threshold (= 7.1 m/s) (Fig.5a) suggesting that drifting snow near the surface was 145 present and after sublimation contributed to the observed sodium peak. The increase in SSA num-146 ber densities and atmospheric sodium occurred in the afternoon a few hours after wind speed had 147 dropped again below the threshold, consistent with a similar phasing observed during the blowing 148 snow event on 14-16 July and discussed in section 3.2.2. To better illustrate episodes of snow drift 149 we include in Figures 4-6 of the revised manuscript a horizontal line marking the estimated threshold 150 wind speed U_t . 151

RC2: Frey et al present an observational study of sea salt aerosol (SSA) production from blow-153 ing snow above sea ice, through measurements during winter 2013 in the Weddell Sea, Antarctica. 154 Since the modelling hypothesis presented by Yang et al (2008, GRL), the mechanism of SSA produc-155 tion from blowing snow has been implemented in numerous modelling studies, unfortunately without 156 observational evidence of the mechanism itself. This work provides a detailed study of the pro-157 posed mechanism through measurements of size distributions and inorganic chemical composition 158 of aerosols and blowing snow, and comparisons to modelled parameters of blowing snow SSA pro-159 duction. Given the prevalence of the use of the blowing snow SSA production parameterisation, this 160 is a very valuable study. 161 **Reply:** We thank the reviewer for the positive assessment. 162 163 **RC2:** My comments mainly focus on clarification of the manuscript and assessment of statistical 164 significance throughout. Given the significant length and many figures and tables, the authors are 165 encouraged to consider moving some material to a supplementary information file if appropriate. 166 **Reply:** We consider tables and figures all essential, in agreement with reviewer 1. However, addi-167 tional material as suggested is now included and presented in a supplement (see below). 168 169 **RC2:** One overarching and major comment that needs to be addressed throughout the manuscript 170 is for uncertainties (or standard deviations) to be listed with average values. This is important for as-171 sessing data variability, as well as for assessment of statistical significance. Indeed, statistical tests 172 of significance should be applied to inform whether 'trends' and 'differences' are indeed statistically 173 significant, which would greatly strengthen the findings presented in the manuscript. This is impor-174 tant because trends sometimes seem to be overstated in the text when compared to large scatter

tant because trends sometimes seem to be overstated in the text when compared to large scatter
 shown in the figures. Routine statements of statistical significance would significantly strengthen the
 conclusions throughout.

Reply: We added standard deviations for all averages reported in the tables (Tables 1, 2) and include in the final manuscript significance of trends and differences, where appropriate. As an example below updated text in section 3.4.4.

¹⁸¹ During storms median atmospheric sea salt concentrations from both estimates showed increases ¹⁸² above background values (Fig. 17a) that were statistically significant based on the Wilcoxon rank-¹⁸³ sum test (p<0.01).

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RC2: *I highly recommend reorganising the manuscript to improve readability. Section 3.2 relies significantly on depletion factors. Therefore, I recommend reorganising to move Sections 3.4.1-3.4.3 to*

¹⁸⁷ be before Section 3.2. Also, the current Section 3.4.3 would be best after Section 3.3.

Reply: We are considering this suggestion for the final manuscript.

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RC2: *Major Comments: Page 1, Line 21 & Page 25, Lines 14-15: These sentences state generally that 'similar processes take place in the Arctic', yet no supporting discussion is provided. Since the current work focuses on the specific conditions of the Antarctic work and no data are provided to evaluate this statement, these sentences should be removed.*

Reply: Agreed. We removed the sentence referring to the Arctic from abstract and conclusions, and
 added text in the conclusions as follows:

Similar in situ measurements are needed to corroborate the importance of sea salt aerosol produc tion from blowing snow also in the Arctic to validate atmospheric and ice core models (e.g. Rhodes
 et al., 2017; Huang and Jaeglé, 2017).

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RC2: Page 1, Lines 2-3 and Page 3, Lines 5-7: The statement 'validating a model hypothesis to account for winter time SSA maxima in polar regions not explained otherwise' generalises beyond the Antarctic, which is not appropriate, and it also not consider other factors, such as lower boundary layer height and lead-based SSA production. This statement should be rephrased to focus on validating wintertime SSA production from blowing snow (which is excellent), as a comprehensive discussion of wintertime SSA maxima causes in both the Arctic and Antarctic is not presented in this work. Further, the work of Huang and Jaegle (2017) did not consider the observed influence of lead-based SSA production in the Arctic (May et al. 2016, JGR). I suggest focusing on the Antarctic,

²⁰⁸ as this is the strength of this work.

Reply: We agree and rephrased in abstract (1.) and introduction (2.) accordingly.

Two consecutive cruises in the Weddell Sea, Antarctica, in winter 2013 provided the first direct observations of sea salt aerosol (SSA) production from blowing snow above sea ice, thereby validating a model hypothesis to account for winter time SSA maxima in the Antarctic.
 Indeed, model agreement with SSA winter maxima observed at a number of locations in the polar regions is much improved when a SSA source from blowing snow based on the parameterisation of (Yang et al., 2008) is included in the model (Huang and Jaeglé, 2017; Yang et al., 2019).

RC2: Figure 1; Page 3, Lines 30-33; Page 13, Lines 22-23: Please provide a legend for sea ice concentration. It appears that stations S2, S3, and S9 were in areas of reduced sea ice concentration. While there is significant evidence for blowing snow SSA production based on chemical analyses, a discussion of the distance to open leads, in addition to open water (Page 3, Line 32), needs to be included, since there is measurement evidence of wind-dependent lead-based SSA production (e.g., Nilsson et al. 2001, JGR).

Reply: A legend is now included in Fig. 1 and we added the following text in section 2 (1.) and in the discussion p13 - after line26 (2.).

Sea ice concentrations in mid July 2013 derived from Nimbus-7 satellite microwave radiometer
 measurements (Comiso, 2018) show areas with 85-95% ice cover near ice stations S2-3 and S7-9
 indicating that open leads may be present (Figure 1).
 Open leads, which may have been present
 in areas of reduced sea ice concentration e.g. near ice stations S2-3 and S7-9 (Figure 1) are another
 potential wind-dependent source of SSA from open water as observed in the Arctic (Nilsson et al.,

2001; May et al., 2016), albeit with a much smaller flux contribution per surface area compared to the 231 open ocean due to reduced fetch and low fraction of surface coverage (<15%).

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RC2: *Page 7, Lines 3-5: Please clarify whether these time periods of ship exhaust influence were also removed from the aerosol size distribution data, as they should be.*

Reply: Ship exhaust influence on measurements of aerosol size and concentration was removed by using a wind-sector filter (section 2.2). We clarified text in section 2.2 including also the fraction of CLASP data filtered out as follows (see corresponding reply to BC1):

²³⁷ CLASP data filtered out as follows (see corresponding reply to RC1):

Raw aerosol number concentrations at the crow's nest showed significant spikes, when air came from
the direction of the ship's engine stack, whereas no evidence of pollution was detected in the observations on the sea ice. Pollution spikes were effectively filtered out prior to averaging by excluding all
data when relative wind direction was in the 135–225° sector encompassing the ship's engine stack.

A total of 21% of the available 1-second data was removed from the crow's nest data.

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RC2: Page 7, Line 9 and Table 3: LODs are normally defined as 3*sigma, rather than 2*sigma. What is the authors justification here? Also, LODs should be reported with one significant figure (too many shown in Table 3, which can be misleading).

Reply: Here we follow Wagenbach et al. (1998) who employed a similar aerosol filter method and de-

fined the mean detection limits as 2*sigma. Two figures for LOD were reported in Table 3 to account
for increased LOD at the shorter run times of filters deployed on the sea ice. However, we removed
that line to report only one figure for LOD and clarified the footnote of Table 3 as follows.

²⁵¹ ^cbased on crow's nest mean air sample STP-volume (6.4 m³); mean air sample STP-volume for filters ²⁵² deployed on the sea ice was 3.3 m³ increasing respective LODs by a factor 1.6

RC2: Tables 4-5: Data below the LOD should be labeled as such, as exact values below LOQs are not meaningful.

Reply: Agreed. Snow concentrations (Tab. 2) were typically 2 orders of magnitude above the LOD of $\sim 2 \text{ ng g}^{-1}$, whereas some aerosol concentrations (Tab. 1) where below the estimated LOD. In the final manuscript a corresponding footnote is added to those values in Table 4:

²⁵⁸ final manuscript a corresponding footnote is added to those values in Table 4:

²⁵⁹ ^cbelow the estimated LOD (see Table 3)

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RC2: Page 8, Lines 3-5: Instead of reporting depletion factors, I highly encourage the authors to consider reporting 'enrichment factors' (e.g. Krvanek et al. 2012, Atmos. Environ.), which are more intuitive to understand in my opinion (i.e. enrichments are >1, depletion corresponds to <1).

Reply: Deviations from bulk sea water ion rations are reported in the literature in both ways, either as enrichment or as depletion factors (e.g. Sander et al., 2003). Since the focus here is on depletion processes we choose to report depletion rather than enrichment factors, also to be consistent with some of the previous related work (Yang et al., 2008). To help interpretation we added a sentence to section 2.4:

For example, $DF_x = -1.5$ or 150% enrichment means the respective ion concentration is 2.5 times that in reference sea water.

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RC2: Page 8, Lines 8-11: I am quite concerned that data were selectively removed from the datasets 272 presented. I can understand if certain samples are not used for externally identified reasons, but if, 273 for example, sulfate concentration is removed for a given sample, I'm concerned about continuing 274 to use other ions from that sample, as appears to have been done based on the numbers shown in 275 Tables 4 and 5. I worry that the presented datasets are skewed based on the removal of these data 276 points. What fraction of the time did ship emissions impact the dataset? It needs to be clarified what 277 fraction of the data were removed. This data treatment is very important for later statements about 278 the distribution of depletion factors (e.g., statements on Page 10, Lines 7-9). 279

Reply: No snow data were removed whereas the fraction of aerosol filter data removed was relatively
 small, and is now mentioned in the revised text. Filter samples suspected of contamination based
 on anomalous sulfate enrichment (total of 6 samples) are not anymore used in the statistics. The
 pollution impact on filter chemistry is now discussed (see reply to RC1 above). Bromide depletion
 factors below a threshold of -7 are considered outliers and removed. The corresponding statistics in
 Table 4 are updated. Follow up statements are not affected by any of these changes.

A total of 6 (= 6% of all crow's nest samples) $DF_{SO_4^{2-}}$ values were below that of pure mirabilite (= -7.3) and are attributed either to sulfate contamination from the ship's engine emissions discussed below or measurement error. We therefore removed all ion concentrations of the corresponding filter samples from the dataset. $DF_{Br^{-}}$ only below -7 were considered outliers due to measurement error and removed: a total of 4 (= 3% of all samples) from the crow's nest data, and a total of 6 (= 14% of all samples) from the sea ice data.

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RC2: *Page 9, Lines 28-30; Page 10, Lines 1-3: Please reference where these data are presented, or please add them as supplementary information.*

Reply: Agreed. In a supplement we include now a Figure S1 (Fig. 2) with an overview of the available observations during ANT-XXIX/7, and Table S1 and Table S2 (Table 3, 4) with the statistics of particle

²⁹⁷ concentration and size. The text has been amended as follows:

1. At 29 m mean total number densities N_{46-478} were 8.7×10³ m⁻³ during ANT-XXIX/6 and very similar 7.2×10³ m⁻³ during ANT-XXIX/7 (Table S1, Figure S1c).

2. At 29 m mean total number densities $N_{0.4-12}$ were 2.1×10⁶ m⁻³ during ANT-XXIX/6 (Table S2, Figure 2d). $N_{0.4-12}$ mean values at 2.0 and 0.2 m during ice stations were 1.4×10⁶ and 1.7×10⁶ m⁻³, respectively, about the same as the number densities observed during the same time at 29 m (Table S2). The median aerosol particle diameters $\overline{d_p}$ at the measurement heights 0.2, 2.0m and 29 m

ranged between 0.60 and 0.66 μ m (Table S2) showing dominance of sub-micron sized particles in atmospheric aerosol below the instrument particle size cut-off (>11 μ m).

306 3. Median $DF_{SO_4^{2-}}$ values at 29 m were very similar during ANT-XXIX/6 (=0.34) and ANT-XXIX/7 307 (=0.30), but larger near the sea ice surface (=0.49), suggesting throughout a significant contribution 308 to the total SSA burden from a fractionated sea ice source (Table 4, Figure S1e).

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RC2: Section 3.4.2 and associated text in Conclusions: The authors should be mindful that only aerosol and snow bromine were measured and that no measurements of reactive bromine are pre-

³¹² sented. Therefore, the strength of the implications for reactive bromine production should be weak-

ened to account for this uncertainty and other factors that contribution to reactive bromine production and abundance.

Reply: Indeed, we do not infer any details on speciation of reactive bromine chemistry. We added a sentence in section 3.4.2 (1.) and amended a sentence in conclusions (2.):

1. Detailed measurements of participating bromine species in air, snow and aerosol are needed to
 further understand relevant processes and constrain the mass budget.

2. It is found that SSA produced by blowing snow is depleted in bromide suggesting it is a source of reactive bromine to the atmosphere, which then can contribute to ozone depletion events.

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RC2: Page 21, Lines 9-10: Depletion factors examine the degree of depletion, but they do not provide information on the mass present. Therefore, the data here cannot assess contribution to the fraction of net bromine release, as currently presented, especially without reactive bromine measurements.

Reply: Agreed, we don't discuss detail of the bromine mass budget. We amended the corresponding sentence in section 3.4.2, conclusions and abstract (1.), as well added a note (2.) (see reply to previous comment):

previous comment):
1. On average snow on sea ice and blowing snow showed no or small depletion of bromide relative
to sodium with respect to sea water, whereas aerosol at 29 m was depleted suggesting that significant bromine loss takes place in the aerosol phase between 2 and 29 m above the sea ice surface.

2. Detailed measurements of participating bromine species in air, snow and aerosol are needed to

³³² further understand relevant processes and constrain the mass budget.

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RC2: Page 19, Lines 22-25: This analysis is only valid if you assume there is no precipitation of NaCI.2H2O. Please verify that based on temperature, and perhaps take out the very low temperature points.

Reply: A complete model of freezing seawater is beyond the scope of this study. Thus we acknowl edge that precipitation of NaCI.2H2O introduces some uncertainty to this analysis by adding the
 sentence below.

Further Na⁺ depletion may arise from the precipitation of hydrohalite (NaCl·2 H₂O) once ambient temperature drops below the threshold of -22.9 °C (e.g. Butler et al., 2016), which occurred here during some periods of time (Fig.2b). In the analysis below however we consider only the precipitation of mirabilite.

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RC2: Page 19, Lines 127-28: Does this also mean that the aerosols collected were a mixture of sea
 salt emitted from the ocean and sublimation of blowing snow?

Reply: Mixing with a pool of non-fractionated sea salt aerosol from the open ocean (DF_{Na^+} , $DF_{SO_4^{2-}}=0$) would move data points towards the origin in Figure 14, but would not explain apparent Na⁺ enrichment or Cl⁻ loss in aerosol at a given SO₄²⁻ depletion. We believe a plausible explanation for the deviation of aerosol observations from the mirabilite precipitation model is HNO₃ induced Cl⁻ loss

from sea salt either in airborne SSA or as an artefact on filters, as stated in the reply to reviewer 1.
 Below we repeat the amended text.

Snow on sea ice follows closely the theoretical mirabilite fractionation line, whereas aerosol shows 353 large scatter and a tendency to apparent Na⁺ enrichment with respect to Cl⁻ of up to 20%, equivalent 354 to Cl⁻ depletion with respect to Na⁺ of 17% (Fig. 14). Dechlorination of sea salt aerosol observed 355 in Antarctica has a maximum in spring/summer, when gaseous acidic species (nitric, sulfuric and 356 methanesulfonic acid) are available to replace chloride on sea-salt aerosol (Wagenbach et al., 1998; 357 Rankin and Wolff, 2003; Legrand et al., 2017). Acidic sulphur species are close to zero during winter 358 in coastal Antarctica e.g. at Neumayer (Weller et al., 2011), whereas nitric acid is low but non-zero, 359 e.g. 1-2 pptv at Halley (Jones et al., 2011). Thus nitric acid induced CI⁻ loss from sea salt is a plausi-360 ble explanation for the observed CI⁻ depletion either in airborne SSA or as a sampling artefact from 361 sea salt already accumulated on the filter surface as suggested previously (Wagenbach et al., 1998; 362 Legrand et al., 2017). Unfortunately no usable filter data of aerosol nitrate are available from this 363 study to further test the association between nitrate and sea salt due to a very high lab procedure 364

365 blank.

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RC2: Page 22, Lines 32-33: A conversion factor is used to calculate [SSA] based on Na+ and using seawater composition, but this seems to undermine and not take into account the sulfate-depletion observed.

Reply: It does not. The impact of the depletion due to mirabilite precipitation on our calculation is indeed very small, and is therefore neglected. We added the text below to clarify:

As shown in section 3.4.1 depletion of SO_4^{2-} due to the precipitation of mirabilite decreases Na⁺ by up to 12%. Reduction in both ions decreases the mass fraction of Na⁺ in the depleted sea salt aerosol by a maximum of ~0.7% compared to reference seawater. Thus, by not considering the depletion effect conversion factor and calculated SSA mass are underestimated by up to ~0.7%, which is negligible given all other uncertainties.

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RC2: Page 21, Lines 30-31 and elsewhere: Is this U10m and the associated data in Fig 16 an average, or threshold? It isn't clear how the data were binned. Please clarify calm and stormy conditions. Does calm represents U10m<5 m/s? How about stormy?

Reply: We used a relatively narrow wind speed range for calm and windy conditions. We amended this to include more data, particularly for the open ocean case when only a few days of measurements were available. Aerosol data are now selected based on a wind speed threshold: calm conditions when $U_{10m} < 4 \text{ m s}^{-1}$ and windy conditions when $U_{10m} > 9 \text{ m s}^{-1}$. We updated Figure 16, including also the standard deviation of the mean, to show statistical significance of differences in size distributions, as suggested further below (Fig. 3). And the text in section 3.4.3 is clarified as follows:

Average aerosol number density and volume distributions observed in the Weddell sea show that 387 during calm conditions (U_{10m} <4 m s⁻¹) concentrations across most of the size spectrum were smaller 388 above sea ice than above the open ocean (Fig. 16a). Depending on particle size the variability was 389 relatively large as illustrated by the standard deviation of the mean values (Fig. 16a). Thus differ-390 ences in mean size distributions were statistically significant only for $d_p < 2 \mu m$ in the case of aerosol 391 number density, and d_p 1-8 μ m in the case of aerosol volume distributions (Fig. 16b). The wind speed 392 threshold chosen for calm conditions is well below the mean snowdrift threshold wind speed U_t of 393 7.1 m s⁻¹ observed during this study and within the range when breaking of waves commences (3-394 4 m s⁻¹; O'Dowd et al., 1997). ... During stormy conditions ($U_{10m} > 9 \text{ m s}^{-1}$) average aerosol number 395 densities above sea ice increased significantly for particle diameters $d_p < 2 \mu m$, reaching at the lower 396 end of the size spectrum levels similar to those observed above the open ocean (Fig. 16a). Average 397 aerosol volume concentrations above sea ice also showed an increase during storms, significant for 398 particle sizes d_p 0.8 to 9 μ m (Fig. 16b). 399

400

RC2: Page 22, Lines 16-19: It seems "not all water is lost" could represent a large uncertainty of blowing snow sublimation. This is important for reactions that depend on the surface area of aerosols. It could be highlighted in the abstract or conclusion. Also, please justify how to get $10^{-3} \mu m$. Using snow salinity of 0.06 psu from Table 5, median snow particle of 100 um from Table 6, yields d(dry) of 1 um.

Reply: We agree the degree to which water ice is lost on particles during sublimation has implica tions for heterogeneous chemistry, something future experiments will need to address; text below has
 been added to the conclusions.

The degree of water ice loss from particles has implications for particle surface area and heterogeneous chemistry, which future experiments will need to address.

Reply: We disagree regarding the calculation of d_{dry} : to convert S_p from psu (equivalent to g of dissolved salt per kg of sea water as defined in Section 2.4) into units of kg per kg in order to be consistent with units of density (kg of salt per m³ of salt) requires division by one thousand as the equation states (Page 22, Line 16), correctly yielding d_{dry} of ~1 nm.

415

416 RC2: Page 23, Lines 1-3: Please show this comparison and data in a supplementary file.

⁴¹⁷ **Reply:** We included Figure S3 in the supplement to show the comparison (Fig. 4), and amended the

418 sentence as follows:

The sea salt mass estimates show that most filter-based values have a low bias compared to median sea salt concentrations derived from $N_{0.4-12}$ during filter sampling intervals (Fig. S3), on average of ~26%. The bias shows also a weak but significant positive correlation wind speed (R=0.4, p<0.01) (Fig. S3). A low bias of the filter samples especially during high wind speeds is expected because the smaller cut-off diameter (<6 μ m) compared to the optical particle counter (>11 μ m) limits capture of coarse sea salt aerosol, where much of the particle mass is located (Fig. 16b).

RC2: Page 25, Lines 5-10: This is not a new finding and has been presented in other work. Therefore, either these sentences should be removed here or other work should be referenced to further support these findings.

Reply: Presenting the links between snow salinity, differences in sea ice age and SSA source strength of blowing snow together with direct observations is of course new. However, we reference relevant work on sea ice and snow on sea ice as follows:

- at a given salt migration distance from the sea ice surface it is total snowpack depth, that determines
the salinity probability distribution of snow on sea ice consistent with previous studies (Domine et al.,
2004; Massom et al., 2001). FYI can therefore be distinguished from MYI based on snow salinity, because snow on FYI is in general more shallow than on MYI. Secondary factors potentially increasing
the difference in salinity between FYI and MYI and identified previously (e.g. Massom et al., 2001)

are more frequent flooding of FYI with seawater due to negative freeboard and MYI desalination due
 to brine drainage.

439

RC2: Data Availability: Since the current work is expect to be very valuable for informing future mod-

elling work and other studies, I highly encourage the authors to put these data in a public archive.

Reply: All data from this study used are stored in the UK Polar Data Centre. The DOI is provided in
 the final manuscript.

All data are stored in the UK Polar Data Centre, Natural Environment Research Council, UK Research
 and Innovation (https://doi.org/10.5285/853dd176-bc7a-48d4-a6be-33bcc0f17eeb, Frey et al., 2019).

RC2: Figure 7: Please add a legend to give meaning to the colors presented. Also, it is stated throughout the manuscript that the surface snow is typically significantly sulfate depleted (justifying the sea ice source for sulfate-depleted aerosol), but here the surface is more often near 0. Please clarify.

Reply: Figure 7 now includes a legend (Fig. 1). There is significant spatial heterogeneity in the sampled local snowpack profiles, whereas blowing snow integrates over a wider area of sea ice. We clarified the discussion of the snow pit observations (Page 13 - Lines 9-14) as follows:

DF_{SO²⁻} profiles exhibited large scatter: except at one location surface-near snow showed no or small 454 depletion, whereas most profiles showed significant depletion in deeper layers within 5-10 cm of the 455 sea ice surface (Fig. 7c). ... However, the $DF_{SO_4^{2-}}$ values of blowing snow were at the top end of the 456 range observed only in the deeper and more saline local snowpack (Fig. 6c). A plausible explanation 457 for this observation during the storm on 15 July is that blowing snow integrates snow contributions 458 from a wider area. And given the spatial heterogeneity of local snowpack thickness and composition 459 blowing snow contributions must have dominated from areas where fractionated snow was at or near 460 the surface such as seen in one of the profiles sampled on 12 July (Fig. 7c). 461

462

RC2: The highly relevant work of Giordano et al. (2018, ACP) 'The importance of blowing snow to halogen-containing aerosol in coastal Antarctica: influence of source region versus wind speed' should be considered in this manuscript.

⁴⁶⁶ **Reply:** We agree and correct the oversight by referring to this work in the introduction:

⁴⁶⁷ A recent observational study in the Ross Sea sector of coastal Antarctica also shows a significant ⁴⁶⁸ association between increased SSA and high wind speed suggesting a link to blowing snow above ⁴⁶⁹ sea ice as a source (Giordano et al., 2018).

- 471 RC2: Minor/Technical Comments: Throughout the manuscript, watch for 'paragraphs' that are only
- ⁴⁷² 1-2 sentences, as this disrupts the flow and limits discussion. Consider reorganization to prevent this.
- ⁴⁷³ **Reply:** We reorganised, where appropriate.
- 474
- 475 **RC2:** Page 1, Line 9: Please state the size of the sulphate-depleted aerosol.
- 476 **Reply:** Done.
- 477 Similar depletion in bulk aerosol observed in the 1-6 μ m range suggests that most sea salt originated
- from snow on sea ice and not the open ocean or leads, e.g. on average \sim 93% during the 8 June and 12 August 2013 period.
- 480
- **RC2:** Page 1, Line 13: Based on the data presented later, 'enriched' is likely a typo and should be 'depleted' here with respect to aerosol at 29 m.
- ⁴⁸³ **Reply:** This is now corrected (see reply to reviewer 1 above).
- 484

488

- **RC2:** *Page 2, Line 20: Provide a reference to a SSA review here.*
- ⁴⁸⁶ **Reply:** The reference below is now included.
- 487 de Leeuw et al. (2011)
- **RC2:** Page 4, Lines 27-28: I think it is dividing kappa instead of multiplying. Please check. Also, please provide the value for the von Karman constant in parentheses.
- ⁴⁹¹ **Reply:** Corrected as follows:
- ⁴⁹² To do this a logarithmic wind profile U(z) is assumed given by $U(z) = u_*/\kappa \ln(z/z_0)$ (e.g. Li and Pomeroy,
- ⁴⁹³ 1997), with measurement height *z*, the von Karman constant κ (= 0.4), friction velocity u_* and the ⁴⁹⁴ surface roughness length of momentum z_0 set to 5.6×10⁻⁵ m as measured very consistently above
- snow at Halley (King and Anderson, 1994).
- 496

RC2: Page 5, Lines 12-14: Please provide a greater description of the inlet. Also, please clarify whether the data presented where corrected for these particle loss estimates ('we adopt' is confusing phrasing).

- 500 **Reply:** Clarified as follows.
- ⁵⁰¹ Particle losses to inlet walls are minimised by using a short and straight inlet tube of 0.3 m length sim-
- ilar to the original configuration (Hill et al., 2008, Figure 9). We assume as an upper limit of particle
 losses those estimated previously for a similar inlet configuration (Norris et al., 2012), which amount
- to 43% at $d_p = 11.32 \,\mu\text{m}$, 19% at $d_p = 6.06 \,\mu\text{m}$ and 0.1% at $d_p = 0.44 \,\mu\text{m}$, respectively.
- **RC2:** Page 5, Lines 22 and 27: Please clarify the size range of aerosol collected.
- ⁵⁰⁷ **Reply:** Clarified as follows.
- ⁵⁰⁸ Filters were estimated to collect aerosol in the diameter range $\sim 0.3 \,\mu$ m to less than $6 \,\mu$ m. The lower ⁵⁰⁹ end of the range is based on previous measurements of collection efficiencies of PTFE filters as a
- function of particle size (Soo et al., 2016), whereas the upper end is based on the estimated cut-off diameter described below.
- 512

- **RC2:** *Page 9, Line 8: I assume the authors are discussion temperature in degrees Celsius, but this needs to be stated.*
- 515 **Reply:** Added.
- 516 Near-zero or positive ambient temperatures T_a in degrees Celsius ...
- 517
- **RC2:** *Page 9, Line 14: Where is the timing of the snowfall presented/shown?*
- **Reply:** Only the timing of airborne snow particles is shown. Occurrence of precipitation is based on 3-hourly ship's weather reports and occasional webcam images. We rephrased accordingly.
- ⁵²⁰ 3-hourly ship's weather reports and occasional webcam images. We rephrased accordingly. ⁵²¹ Winter storms occurred frequently with wind speeds ranging between 10 and 20 m/s, occasionally
- exceeding 20 m/s, and coincided often with snowfall based on the ship's 3-hourly weather report,
- ⁵²³ occasional webcam images and presence of clouds (data not shown).

524

- **RC2:** Page 9, Line 22-23: Please provide a reference that connects the friction velocity with the boundary layer conditions. Also, reference where these data are shown, or add to a SI.
- **Reply:** We included Jacobson (2005) and Nishimura and Nemoto (2005) as references, as well as a figure (Fig. 5) in the supplement showing the correlation between friction velocity u_* and horizontal wind speed U.

530

- **RC2:** Page 10, Line 16: Please clarify 'two 7-10 day long periods'. I'd suggest wording such as 'two periods, one lasting 7 days and another 10 days', or similar.
- 533 **Reply:** Added as follows.
- Two periods, one lasting 7 days and another 10 days, were chosen based on data coverage to discuss key features of observed blowing snow and associated SSA increases.
- 536
- **RC2:** Page 11, Lines 3-4: Please provide concentrations in parentheses for context.
- 538 **Reply:** Clarified as follows.
- ⁵³⁹ Near the surface spectral number densities N_{0.4-12} for particles with $d_p < 2 \mu m$ during the storm on 24
- June remained with 10^5 m^{-3} below those seen at 29 m (10^6 m^{-3}) likely due to scavenging of aerosol by snow particles (Fig. 4d-e).
- 542
- **FC2:** Page 13, Lines 16-17: The direct comparison of $N_{0.4-12}$ to $d_p p < 2\mu m$ here is confusing since
- these are different size ranges.
- 545 **Reply:** Clarified as follows.
- Aerosol size spectra show that number densities of particles with size $d_p < 2 \mu m$ increased during individual storms by 2-3 orders of magnitude above background levels.
- 548
- **RC2:** *Page 14, Line 15: Please define SWE (snow water equivalent?) and the 'saltation layer' (what height?).*
- 551 **Reply:** Amended as follows.
- 1. (mm day⁻¹ snow water equivalent) 2. The saltation layer is a layer just above the snow surface
 usually several centimetres thick (e.g. Déry and Yau, 1999).
- 554

- **RC2:** Page 15, Line 3: What does '(0.001)' correspond to here? Please clarify.
- 556 **Reply:** Amended as follows.
- ⁵⁵⁷ ... when snow drift density μ right above the snow surface exceeds a critical value μ_c (= 0.005 kg ⁵⁵⁸ m⁻³). For comparison a lower value of μ_c (= 0.001 kg m⁻³) is also considered.
- **RC2:** Page 15, Lines 6 and 11: Please clarify that Ut and u*t are calculated, not observed.
- **Reply:** U_t and u^{*} are not calculated. Windspeed and snow particle number densities are both measured quantities; thus drift threshold wind speed is an observed quantity based on the combination of two measurements (symbols in Figure 8) as opposed to modelled values (Eq 4). Similar for friction velocity u^{*}. We added a sentence to clarify.
- The observed threshold wind speed U_t and friction velocity u_t^* are the respective measurements at the onset of drifting or blowing snow.
- 567
- **RC2:** Page 15, Line 15: Please show how u*t values were calculated.
- **Reply:** u^{*} is not calculated. See reply above.
- 570
- **RC2:** Page 15, Line 32: Please define what you mean by 'minor' here. Please quantify.
- **Reply:** We did not run the model but the model bias in absolute values will cancel out because ratios
 are used (see Eq.2). Clarified as follows.
- The model bias in q_{bsalt} is expected to cancel out in estimates of bulk sublimation rate Q_s (Eq. 2) and
- therefore also of SSA production Q_{SSA} (Eq. 1) because the calculation uses not absolute values but ratios of actual q_{Last} and its maximum q_{Last}
- ratios of actual q_{balt} and its maximum q_{b0} .

- 577
- 578 RC2: Page 17, Line 13: Please delete 'have' typo.
- 579 **Reply:** done
- 580
- **RC2:** Page 17, Line 3: Didn't mean d_p increase?
- **Reply:** Decrease is correct. Expected is a decrease of d_p with height above the surface snow particle source in the absence of snowfall due to gravitational settling.
- 584
- ⁵⁸⁵ **RC2:** Page 19, Line 32: Do you mean 0.1204 here?
- **Reply:** This has been corrected.
- 587
- **RC2:** Page 20, Lines 8-10: The wording 'well established' should be removed, as the Yang et al papers are models based on a hypothesis rather than measurement based and this associated uncertainty should be noted.
- ⁵⁹¹ **Reply:** Agreed and clarified as follows.
- ⁵⁹² Modelling studies suggest that sea salt may be an important source of atmospheric bromine species
- in the mid to high southern latitudes, and that SSA from blowing snow releases bromine (Yang et al.,
- ⁵⁹⁴ 2008, 2010) driving ozone depletion events observed during or after snow storms (Jones et al., 2009).
- 595
- **RC2:** Page 20, Line 27: Data in Table 5 are presented in $\mu g g^{-1}$. Please fix or clarify.
- ⁵⁹⁷ **Reply:** Corrected as follows.
- ⁵⁹⁸ Median bromide concentrations in snow ranged between 0.07 and 0.18 μ g g⁻¹ (Table 5).
- 599

- **RC2:** Page 21, Line 11: Change "due a" to "due to a".
- 601 **Reply:** Corrected.
- **RC2:** Page 21, Line 14: No data are presented examining the acidity of the surface snowpack.
- **Reply:** Agreed, pH of aerosol and snow was not measured. We therefore removed reference to acidity.
- The bromine release from SSA produced by blowing snow may be more efficient because it has a large fraction of sub-micron sized particles (see section 3.4.3), and resides at the well ventilated top of the blowing snow layer.
- 609
- RC2: Page 23, Line 22: Delete extra "the".
- 611 Reply: Corrected.
- 612
- **RC2:** Page 23, Lines 29-30: Remove "always" and replace with "often" to more appropriately reflect
- 614 the data shown.
- 615 **Reply:** Agreed and amended.
- 616
- 617 **RC2:** Page 25, Line 27: "LL & MM"?
- 618 **Reply:** Mentors who prefer to remain anonymous
- 619
- **RC2:** Figure 16: The variations in these distributions (e.g. standard deviations) should be shown.
- **Reply:** Agreed, we updated Figure 16 including the standard deviation of the mean values, and corrected the caption (Fig. 3).
- 623
- **RC2:** Figure 17: This figure is difficult to understand currently.
- ⁶²⁵ **Reply:** We updated Figure 17 and clarified the caption (Fig. 6).
- 626

SC1: This manuscript describes an interesting set of measurements and detailed analysis confirm-627 ing the blowing snow as a significant source for sea salt aerosol in the vicinity of sea ice in coastal 628 Antarctica. We agree that this is an important result with significant implications for polar tropospheric 629 aerosol loadings and heterogeneous halogen chemistry. However, it would be helpful to both the au-630 thors and readers of this article to refer to prior work also published in ACP showing similar results 631 from measurements taken on sea ice in the Ross Sea. Giordano et al., 2018 also clearly identifies 632 blowing snow on sea ice as a significant source of chlorine rich sea salt aerosol from online Aerosol 633 Mass Spectrometer measurements of aerosol composition, optical measurements of blowing snow 634 and interstitial aerosol concentrations and offline measurements of surface and blowing snow com-635 position. The consistency between the results from observations using different techniques and on 636 opposite sides of the Antarctic continent further indicates the importance of this mechanism to the 637 overall Antarctic aerosol budget. 638 Lars Kalnajs and Peter DeCarlo 639

640

Reference: Giordano, M. R., Kalnajs, L. E., Goetz, J. D., Avery, A. M., Katz, E., May, N. W., Leemon, A., Mattson, C., Pratt, K. A., and DeCarlo, P. F.: The importance of blowing snow to halogen-

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⁶⁴⁴ Phys., 18, 16689-16711, https://doi.org/10.5194/acp-18- 16689-2018, 201

645

Reply: We agree and apologise for the oversight of this interesting study (see also reply to RC2
 above). We now refer to this work in the introduction:

A recent observational study in the Ross Sea sector of coastal Antarctica also shows a significant
 association between increased SSA and high wind speed suggesting a link to blowing snow above
 sea ice as a source (Giordano et al., 2018).

652 References

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Figure 1: manuscript Figure 7 - Vertical snowpack profiles sampled at various locations on the ice floe of ice station S6 during the 11–14 July 2013 period (color indicates day of sampling): (a) salinity S_p , (b) Na⁺ concentrations and (c) sulfate depletion factor $DF_{SO_4^{2^-}}$ with respect to Na⁺ as a function of snow height above the sea ice surface. Symbols illustrate averages for snow layers of 2 cm thickness, except those with white face color indicating 0.5-1.0 cm layer thickness. Data points at the top of each profile represent the surface snow layer, thus adding half the snow layer thickness to snow height yields total snowpack depth. Shaded areas illustrate the range of the respective parameter measured in blowing snow on 15 July 2013 (Fig. 6e-f).



Figure 2: manuscript Figure S1 - Overview of atmospheric observations in the Weddell Sea from 14 August to 17 October 2013 (ANT-XXIX/7): (a) horizontal wind speed *U* at 39 m. (b) ambient temperature T_a and relative humidity with respect to ice RH_{ice} at 29 m. (c) total number densities N_{46-478} of airborne snow particles at 29 m. (d) aerosol Na⁺ concentrations and (e) sulphate depletion factor $DF_{SO_4^{2^-}}$, both at 29 m.



Figure 3: manuscript Figure 16 - Comparison of mean number distributions (panel a) and volume distributions (panel b) of aerosol above the open ocean (13 to 16 June 2013) and sea ice in the Weddell Sea (18 June to 21 July 2013) during calm (U_{10m} <4 m s⁻¹) and windy (U_{10m} >9 m s⁻¹) conditions. Shaded areas and error bars show the standard deviation of the mean during calm and windy conditions, respectively. Data included are observations from 29 m above the sea surface at ambient *RH*.



Figure 4: manuscript Figure S3 - Comparison of atmospheric sea salt concentrations during the 8 June to 26 July 2013 period derived from filter measurements and from median number densities $N_{0.4-12}$ measured with the CLASP during filter sampling intervals. Data included are observations from 29 m above the sea surface. Symbols are color coded based on wind speed U_{10m} .



Figure 5: manuscript Figure S2 - Comparison between friction velocity u_* and horizontal wind speed U at 2 and 39 m above the sea ice surface. The legends shows respective coefficients of determination of the linear regression. Note that U_{2m} has been derived from the 3-D wind measurements of the sonic anemometer.



Figure 6: manuscript Figure 17 - The partitioning of sea salt between atmosphere and snow, and sulfate depletion above first year sea ice from 18 June to 21 July 2013. Panel (a) shows median atmospheric sea salt concentrations during calm ($U_{10m} < 4 \text{ m s}^{-1}$) and windy ($U_{10m} > 9 \text{ m s}^{-1}$) conditions derived from aerosol filter measurements (filter) and spectral particle number densities $N_{0.4-12}$ (CLASP) (see text). For comparison, a potential atmospheric concentration is calculated assuming that all sea salt observed in the top 0.1 mm of snow on sea ice was released by sublimation and mixed into a 100 m thick boundary layer (for better comparison multiplied here by 0.1). Panel (b) shows for the same time period median sulfate depletion factors $DF_{SO_4^{2-}}$ (with respect to Na⁺) in in surface snow and in aerosol during windy ($U_{10m} > 9 \text{ m s}^{-1}$) conditions. Symbols and errorbars represent median and lower and upper quartiles, respectively.

Table 1: manuscript Table 4 - Descriptive statistics of the aerosol chemistry during ANT-XXIX/6 (ANT6) and ANT-XXIX/7 (ANT7) with mean and median values weighted by the filter sampling interval. Ion and sea salt concentrations are in units of ng m⁻³. See section 2.4 for definition of depletion factors DF.

Parameter	ANT6 at 2 m mean $\pm \sigma$	median	N ^a	at 29 m mean $\pm \sigma$	median	N ^a	ANT7 at 29 m mean $\pm \sigma$	median	N ^a
sea-salt ^b	707 ±1500	336	43	1253 ±2319	639	106	559 ±486	425	28
Na ⁺	217 ±460	103	43	384 ±711	196	106	171 ±149	130	28
CI [_]	379 ± 765	179	43	656 ±1225	302	106	311 ±282	232	27
SO_4^{2-}	28 ^c ±61	19 ^c	38	75 ±152	45	84	33 ±30	23 ^c	28
Br⁻	2.0 ± 1.0	1.9	42	$1.5^{c} \pm 3.0$	0.7 ^c	98	$0.5^{\circ} \pm 0.6$	0.5 ^c	23
$DF_{SO_4^{2-}}$	0.29 ± 0.57	0.48	38	0.07 ± 0.94	0.29	74	0.12 ± 0.60	0.21	27
DF_{Na^+}	-0.08 ±0.29	-0.03	43	-0.46 ±2.29	-0.04	97	-0.02 ±0.19	-0.01	27
$DF_{Br^{-}}$	-1.66 ±1.86	-1.86	36	0.04 ± 0.96	0.37	89	0.05 ± 1.26	0.49	23

^asample size ^bsea salt concentration is derived by multiplying the Na⁺ concentration by 3.262 based on the Na⁺ mass fraction in reference seawater after Millero et al., 2008 ^cbelow the estimated LOD (Table 3) Table 2: manuscript Table 5 - Descriptive statistics of the volume-integrated snow chemistry during ANT-XXIX/6 on first-year sea ice (FYI) at ice stations S1-S6, on multi-year sea ice (MYI) at ice stations S7-9, and for snow layers within 10 cm of the snow surface (TOP10). Ion and sea salt concentrations are in units of μ g g⁻¹. See section 2.4 for definition of depletion factors *DF*.

Parameter	FYI			MYI			TOP10		
	$\mathrm{mean} \ \mathrm{\pm} \sigma$	median	N ^a	$\mathrm{mean} \ \mathrm{\pm} \sigma$	median	N ^a	$\mathrm{mean} \pm \sigma$	median	N^{a}
snow depth (cm)	20.9 ±8.3	19.0	17	50.0 ±32.2	33.0	7	-	-	-
S_p (psu)	1.40 ± 3.99	0.11	110	0.82 ±4.31	0.02	104	0.31 ±0.90	0.06	96
sea salt ^b	1176 ±3518	83	86	590 ±3157	22	95	249 ±729	58	80
Na ⁺	361 ±1079	26	86	181 ±968	7	95	76 ±223	18	80
CI-	680 ±2035	48	87	305 ±1842	13	98	141 ±415	34	81
SO_4^{2-}	61 ±182	6	87	30 ± 166	1	98	17 ±62	3	81
Br⁻	4.28 ±12.23	0.18	85	1.76 ±10.92	0.07	90	1.01 ±3.72	0.12	78
$DF_{SO_{4}^{2-}}$	0.19 ±0.41	0.24	86	0.33 ±0.44	0.35	94	0.27 ±0.39	0.27	80
DF_{Na^+}	0.01 ±0.38	0.06	86	-1.09 ±8.88	0.07	94	-0.11 ±0.99	0.06	80
$DF_{Br^{-}}$	-0.25 ± 0.98	0.05	83	-0.28 ±1.16	-0.01	86	-0.21 ± 0.99	0.04	76

^asample size ^bsea salt concentration is derived by multiplying the Na⁺ concentration by 3.262 based on the Na⁺ mass fraction in reference seawater after Millero et al., 2008

Table 3: manuscript Table S1 - Descriptive statistics of airborne snow particles observed for 8 June to 12 August 2013 (ANT6) and for 14 August to 16 October 2013 (ANT7): total number densities N_{46-478} and particle diameter d_p . Statistics refer to periods when airborne snow particles were present, i.e. times with no snow particles observed were removed prior to averaging.

Parameter	ANT6 at 0.2 m	at 29 m ^a	at 29 m	ANT7 at 29 m
N_{46-478} (m ⁻³)				
mean	2.6×10 ⁵	4.0×10 ³	8.7×10 ³	7.2×10 ³
σ	7.4×10 ³	9.5×10 ³	2.7×10 ⁴	2.2×10 ⁴
median	4.7×10 ³	7.7×10 ²	9.9×10 ²	1.3×10 ³
d_p (µm)				
mean	138	132	133	143
σ	59	59	53	53
median	132	117	124	136
N ^b	8608	11766	42959	37123
sampling time (days) ^c	6	8	30	26

 $^{\rm a}$ for direct comparison of vertical differences statistics of the 29 m measurements only for times when sea ice observations at 0.2 m were available $^{\rm b}$ sample size $^{\rm c}$ total aggregated time during which airborne snow particles were detected

Parameter	at 0.2 m	at 2 m	at 29 m ^a	at 29 m
$N_{0.4-12}$ (m ⁻³)				
mean	1.7×10 ⁶	1.4×10 ⁶	1.4×10 ⁶	2.1×10 ⁶
σ	2.5×10 ⁶	1.9×10 ⁶	1.6×10 ⁶	6.4×10 ⁶
median	8.6×10 ⁵	5.6×10 ⁵	8.0×10 ⁵	1.1×10 ⁶
d_p (µm)				
mean	0.67	0.60	0.67	0.69
σ	0.11	0.06	0.11	0.14
median	0.66	0.60	0.65	0.66
N ^b	13077	14907	9963	48892
sampling time (days) ^c	9	10	7	34

Table 4: manuscript Table S2 - Descriptive statistics of aerosol observed during 8 June - 26 July 2013 (ANT6): total number densities $N_{0.4-12}$ and particle diameter d_p .

 $^{\rm a}$ for direct comparison of vertical differences statistics of the 29 m measurements only for times when sea ice observations at 2 m were available $^{\rm b}$ sample size $^{\rm c}$ total aggregated sampling time