

1 **OBSERVATIONS OF HIGHLY OXIDISED**
2 **MOLECULES AND PARTICLE NUCLEATION**
3 **IN THE ATMOSPHERE OF BEIJING**

4
5 **James Brean¹, Roy M. Harrison^{1*†}, Zongbo Shi¹**
6 **David C.S. Beddows¹, W. Joe F. Acton²**
7 **C. Nicholas Hewitt², Freya A. Squires³ and James Lee³**
8
9

10 **¹Division of Environmental Health and Risk Management,**
11 **School of Geography, Earth and Environmental Sciences**
12 **University of Birmingham**
13 **Edgbaston, Birmingham B15 2TT**
14 **United Kingdom**

15
16 **²Lancaster Environment Centre**
17 **Lancaster University, Lancaster LA1 4YQ**
18 **United Kingdom**

19
20 **³National Centre for Atmospheric Science**
21 **Wolfson Atmospheric Chemistry Laboratory**
22 **University of York, York YO10 5DD**
23 **United Kingdom**
24

* To whom correspondence should be addressed.

Tele: +44 121 414 3494; Fax: +44 121 414 3709; Email: r.m.harrison@bham.ac.uk

†Also at: Department of Environmental Sciences / Center of Excellence in Environmental Studies, King Abdulaziz University, PO Box 80203, Jeddah, 21589, Saudi Arabia

25 **ABSTRACT**

26 Particle nucleation is one of the main sources of atmospheric particulate matter by number, with new
27 particles having great relevance for human health and climate. Highly oxidised multifunctional
28 organic molecules (HOMs) have been recently identified as key constituents in the growth, and,
29 sometimes, in initial formation of new particles. While there have been many studies of HOMs in
30 atmospheric chambers, flow tubes and clean environments, analyses of data from polluted
31 environments are scarce. Here, measurements of HOMs and particle size distributions down to small
32 molecular clusters are presented alongside volatile organic compounds (VOC) and trace gas data from
33 a campaign in June 2017, in Beijing. Many gas phase HOMs have been characterised and their
34 temporal trends and behaviours analysed in the context of new particle formation. The HOMs
35 identified have a comparable degree of oxidation to those seen in other, cleaner, environments, likely
36 due to an interplay between the higher temperatures facilitating rapid hydrogen abstractions and the
37 higher concentrations of NO_x and other $\text{RO}_2\cdot$ terminators ending the autoxidation sequence more
38 rapidly. Our data indicate that alkylbenzenes, monoterpenes, and isoprene are important precursor
39 VOCs for HOMs in Beijing. Many of the C_5 and C_{10} compounds derived from isoprene and
40 monoterpenes have a slightly greater degree of average oxidation state of carbon compared to those
41 from other precursors. Most HOMs except for large dimers have daytime peak concentrations,
42 indicating the importance of $\text{OH}\cdot$ chemistry in the formation of HOMs, as O_3 tends to be lower on
43 days with higher HOM concentrations ; similarly, VOC concentrations are lower on the days with
44 higher HOM concentrations. The daytime peaks of HOMs coincide with the growth of freshly formed
45 new particles, and their initial formation coincides with the peak in sulfuric acid vapours, suggesting
46 that the nucleation process is sulfuric acid-dependent, with HOMs contributing to subsequent particle
47 growth.

48

49

50 1. INTRODUCTION

51 Atmospheric particle nucleation, or the formation of solid or liquid particles from vapour phase
52 precursors is one of the dominant sources of global aerosol by number, with primary emissions
53 typically dominating the mass loadings (Tomasi et al., 2016). New particle formation (NPF) or the
54 secondary formation of fresh particles is a two-step process comprising of initial homogeneous
55 nucleation of thermodynamically stable clusters and their subsequent growth. The rate of growth
56 needs be fast enough to out-compete the loss of these particles by coagulation and condensation
57 processes in order for the new particles to grow, and hence NPF is a function of the competition
58 between source and sink (Gong et al., 2010). New particle formation has been shown to occur
59 across a wide range of environments (Kulmala et al., 2005). The high particle load in urban
60 environments was thought to suppress new particle formation until measurements in the early 2000s
61 (McMurry et al., 2000; Shi et al., 2001; Alam et al., 2003), with frequent occurrences observed even
62 in the most polluted urban centres. NPF events in Beijing occur on about 40% of days annually,
63 with the highest rates in the spring (Wu et al., 2007, 2008; Wang et al., 2016). Chu et al. (2019)
64 review many studies of NPF which have taken place in China and highlight the need for long-term
65 observations and mechanistic studies.

66

67 NPF can lead to the production of cloud condensation nuclei (CCN) (Wiedensohler et al., 2009; Yu
68 and Luo, 2009; Yue et al., 2011; Kerminen et al., 2012) which influences the radiative atmospheric
69 forcing (Penner et al., 2011). A high particle count, such as that caused by nucleation events, has
70 been shown to precede haze events in environments such as Beijing (Guo et al., 2014). These events
71 are detrimental to health and quality of life. The sub-100 nm fraction of particles to which new
72 particle formation contributes to is often referred to as the ultrafine fraction. Ultrafine particles
73 (UFPs) pose risks to human health due to their high number concentration. UFPs exhibit gas-like
74 behaviour and enter all parts of the lung before penetrating into the bloodstream (Miller et al.,
75 2017). They can initiate inflammation via oxidative stress responses, progressing conditions such as

76 atherosclerosis and initiating cardiovascular responses such as hypertension through to myocardial
77 infarction (Delfino et al., 2005; Brook et al., 2010).

78

79 Highly oxidised multifunctional molecules (HOMs), organic molecules with O:C ratios >0.6, are
80 the result of atmospheric autoxidation and have recently been subject to much investigation, in part
81 because the extremely low volatilities arising from their high O:C ratios favour their condensation
82 into the particulate phase. HOMs are most well characterised as the product of oxidation of the
83 biogenic monoterpene compound α -pinene (Riccobono et al., 2014; Tröstl et al., 2016; Bianchi et
84 al., 2017). Although globally, biogenic volatile organic compound (BVOC) concentrations far
85 exceed anthropogenic volatile organic compound (AVOC) concentrations, in the urban environment
86 the anthropogenic fraction is far more significant. Formation of HOMs from aromatic compounds
87 has been demonstrated in laboratory studies and these have been hypothesised to be large drivers of
88 NPF in urban environments (Wang et al., 2017; Molteni et al., 2018; Qi et al., 2018). The formation
89 of HOMs through autoxidation processes begins with the reaction of VOCs with OH, O₃ or NO₃;
90 formation of a peroxy radical (RO₂) is followed by rapid O₂ additions and intra-molecular hydrogen
91 abstractions (Jokinen et al., 2014; Rissanen et al., 2014; Kurtén et al., 2015). Furthermore,
92 generation of oligomers from stabilised Criegee intermediates arising from short chain alkenes has
93 been hypothesised as a contributor of Extremely Low Volatility Organic Compounds (ELVOCs)
94 and Low Volatility Organic Compounds (LVOCs) (Zhao et al., 2015). The low volatilities of these
95 molecules arise from their numerous oxygen-containing functionalities, and this allows them to
96 make a significant contribution to early stage particle growth where other species cannot due to the
97 Kelvin effect (Tröstl et al., 2016), although the contribution of HOMs to the initial molecular
98 clusters is still debated (Kurtén et al., 2016; Elm et al., 2017; Myllys et al., 2017).

99

100 Recent technological advances have facilitated insights into the very first steps of nucleation which
101 were previously unseen, with mass spectrometric techniques such as the Atmospheric Pressure

102 Interface Time of Flight Mass Spectrometer (APi-ToF) and its chemical ionisation counterpart (CI-
103 APi-ToF) allowing for high mass and time resolution measurements of low volatility compounds
104 and molecular clusters. Diethylene glycol based particle counters, such as the Particle Size
105 Magnifier (PSM) allow for measurements of particle size distributions down to the smallest
106 molecular clusters nearing 1 nm. Recent chamber studies have elucidated the contribution of
107 individual species to particle nucleation, ammonia and amines greatly enhancing the rate of sulfuric
108 acid nucleation (Kirkby et al., 2011; Almeida et al., 2013). In these studies, HOMs have been
109 identified, formed through autoxidation mechanisms (Schobesberger et al., 2013; Riccobono et al.,
110 2014; Ehn et al., 2014). These are key to early particle growth (Tröstl et al., 2016) and can nucleate
111 even in the absence of sulfuric acid in chambers (Kirkby et al., 2016) and in the free troposphere
112 (Rose et al., 2018). In this paper, we report the results of HOM and particle size measurements
113 during a summer campaign in Beijing, China.

114

115 **2. DATA AND METHODS**

116 **2.1. Sampling Site**

117 Sampling was performed as part of the Air Pollution and Human Health in a Developing Megacity
118 (APHH-Beijing) campaign, a large international collaborative project examining emissions,
119 processes and health effects of air pollution. For a comprehensive overview of the programme, see
120 Shi et al. (2019). All sampling was conducted across a one month period at the Institute for
121 Atmospheric Physics (IAP), Chinese Academy of Sciences, Beijing (39°58.53'N, 116°22.69'E).
122 The sampling was conducted from a shipping container, with sampling inlets 1-2 metres above
123 ground level, the nearest road being 30 metres away. Meteorological parameters (wind speed, wind
124 direction, relative humidity (RH) and temperature) were measured at the IAP meteorological tower,
125 20 metres away from the sampling site, 30 metres from the nearest road at a height of 120 metres.
126 Data was continuously taken from the CI-API-ToF during a two week period, but due to data losses

127 only five days of data is presented here. Particle size distribution measurements were taken during a
128 33 day period from 24/05/2017 – 26/06/2017.

129

130 **2.2 Chemical Ionisation Atmospheric Pressure Interface Time of Flight Mass** 131 **Spectrometry**

132 The Aerodyne Nitrate Chemical Ionisation Atmospheric Pressure Interface Time of Flight Mass
133 Spectrometer (CI-APi-ToF) was used to make measurements of neutral oxidised organic
134 compounds, sulfuric acid and their molecular clusters at high time resolution with high resolving
135 power. The ionization system charges molecules by adduct formation, such as in the case of organic
136 compounds with two or more hydrogen bond donor groups (Hytinen et al., 2015), or proton
137 transfer in the case of strong acids like sulfuric acid. Hydroxyl or hydroperoxyl functionalities are
138 both common hydrogen bond donating groups, with hydroperoxyl being the more efficient
139 hydrogen bond donor (Møller et al., 2017). This instrument has been explained in great detail
140 elsewhere (Junninen et al., 2010; Jokinen et al., 2012), but briefly the front end consists of a
141 chemical ionisation system where a 10 LPM sample flow is drawn in through the 1 metre length 1”
142 OD stainless steel tubing opening. A secondary flow was run parallel and concentric to this sample
143 flow, rendering the reaction chamber effectively wall-less. A 3 SCCM flow of a carrier gas (N₂) is
144 passed over a reservoir of liquid HNO₃, entraining vapour which is subsequently ionised to NO₃⁻ via
145 an X-ray source. This flow is then guided into the sample flow. . The nitrate ions will then charge
146 molecules either by clustering or proton transfer. The mixed flows travelling at 10 LPM enter the
147 critical orifice at the front end of the instrument at 0.8 LPM and are guided through a series of
148 differentially pumped chambers before reaching the ToF analyser. Two of these chambers contain
149 quadrupoles which can be used to select greater sensitivity for certain mass ranges, and the voltages
150 across each individual chamber can be tuned to maximise sensitivity and resolution for ions of
151 interest. Mass spectra are taken at a frequency of 20 kHz but are recorded at a rate of 1 Hz. All data
152 analysis was carried out in the *Tofware* package in *Igor Pro 6* (Tofwerk AG, Switzerland). A seven

153 point mass calibration was performed for every minute of data, and all data was normalised to
154 signal at 62, 80 and 125 m/Q to account for fluctuations in ion signal, these masses representing
155 NO_3^- , H_2ONO_3^- and $\text{HNO}_3\text{NO}_3^-$ respectively.. Typical values for calibration coefficients range from
156 10^9 - 10^{10} molecules cm^{-3} from these normalised data (Kürten et al., 2012), producing peak sulfuric
157 acid concentrations in the range of 10^6 molecules cm^{-3} . From the very limited periods with
158 simultaneous data for SO_2 , OH radical and condensation sink, it was possible to calculate H_2SO_4
159 concentrations of 10^3 to 10^5 molec cm^{-3} , in which range the calibration constant was $7.0 \pm 1.6 \times 10^8$
160 cm^{-3} which fits well with that expected for this concentration range (Kürten et al., 2012). The
161 nitrate-water cluster is included as the presence of many nitrate-water clusters of the general
162 formula $(\text{H}_2\text{O})_x(\text{HNO}_3)_y\text{NO}_3^-$ were found, where $x = (1, 2, 3 \dots 20)$ and $y = (0, 1)$. No sensitivity
163 calibration was performed for these measurements, and so all values are reported in normalised
164 signal intensity. Due to the high resolving power of the CI-API-ToF system (mass resolution of
165 3500 m/dm and mass accuracy of 20 ppm at 288 m/Q), multiple peaks can be fit at the same unit
166 mass and their molecular formulae assigned. These peaks follow the general formula $\text{C}_x\text{H}_y\text{O}_z\text{N}_w$
167 where $x = 2-20$, $y = 2-32$, $z = 4-16$ and $w = 0-2$, spanning from small organic acids like oxalic and
168 malonic acid through to large dimers of oxidised monoterpene RO_2 radicals such as $\text{C}_{20}\text{H}_{31}\text{O}_9\text{N}$.
169 Beyond 500 m/Q , peak fitting and assignment of compositions becomes problematic as signal
170 decreases, mass accuracy decreases, and the total number of chemical compositions increases, so
171 peaks above the C_{20} region have not been assigned, and a number of peaks have been unassigned
172 due to this uncertainty (Cubison and Jimenez, 2015). As proton transfer mostly happens with acids,
173 and nearly all HOM molecules will be charged by adduct formation it is possible to infer the
174 uncharged formula; therefore all HOMs from here onwards will be listed as their uncharged form.

175

176

177

178

179 **2.3. Size Distribution Measurements**

180 Two Scanning Mobility Particle Sizer (SMPS) instruments measured particle size distributions at
181 15 minute time resolution, one LongSMPS (TSI 3080 EC, 3082 Long DMA, 3775 CPC, TSI, USA)
182 and one NanoSMPS (3082 EC, 3082 Nano DMA, 3776 CPC, TSI, USA) measuring the ranges 14-
183 615 nm and 4-65 nm respectively. A Particle Size Magnifier (A10, Airmodus, FN) linked to a CPC
184 (3775, TSI, USA) measured the sub-3 nm size fraction. The PSM was run in stepping mode,
185 operating at four different saturator flows to vary the lowest size cut-off of particles that it will grow
186 (this cut-off is technically a point of 50% detection efficiency) of <1.30, 1.36, 1.67 and 2.01 nm.
187 The instrument switched between saturator flows per 2.5 minutes, giving a sub-2.01 nm size
188 distribution every 10 minutes. The data was treated with a moving average filter to account for
189 jumps in total particle count, and due to the similar behaviour of the two upper and two lower size
190 cuts, these have been averaged to two size cuts at 1.30 and 1.84 nm.

191

192 **2.4. Calculations**

193 The condensation sink (CS) was calculated from the size distribution data as follows:

$$194 \quad CS = 4\pi D \sum_{d'_p} \beta_{m,d'_p} d'_p N_{d'_p} \quad (1)$$

195

196 where D is the diffusion coefficient of the diffusing vapour (assumed sulfuric acid), β_m is a
197 transition regime correction (Kulmala et al., 2012), d'_p is particle diameter, and $N_{d'_p}$ is the number
198 of particles at diameter d'_p .

199

200 **2.5. Other Measurements**

201 Measurements of the classical air pollutants were measured on the same site, and have been
202 reported in the campaign overview paper (Shi et al., 2019). SO₂ was measured using a 43i SO₂
203 analyser (ThermoFisher Scientific, USA), O₃ with a 49i O₃ analyser (ThermoFisher Scientific,

204 USA) and NO_x with a 42i-TL Trace NO_x analyser (ThermoFisher Scientific, USA), and a T500U
205 CAPS NO₂ analyser (Teledyne API, USA). VOC mixing ratios were measured using a Proton
206 Transfer Reaction-Time of Flight-Mass Spectrometer (PTR-ToF 2000, Ionicon, Austria).

207

208 **3. RESULTS AND DISCUSSION**

209 **3.1. Characteristics of Sampling Period**

210 A total of five days of CI-API-ToF data were collected successfully, from 2017/06/21 midday
211 through 2017/06/26 midday. New particle formation events were observed on 24th June in the late
212 afternoon and 25th June at midday. Some nighttime formation of molecular clusters was seen
213 earlier in the campaign, as were several peaks to the 1.5 – 100 nm size range, likely from pollutant
214 plumes containing freshly nucleating condensable materials. The trace gases, O₃, SO₂, NO and NO₂
215 are plotted in the Figure S1. O₃ shows mid-afternoon peaks, around ~120 ppb on the first two days
216 of the campaign, and 50-70 ppb for the latter days. SO₂ shows a large peak, reaching 4 ppb on 22/06
217 but <1 ppb for the rest of campaign. NO shows strong mid-morning rush hour related peaks,
218 declining towards midday due to being rapidly consumed by O₃. NO₂ shows large traffic related
219 peaks. The sulfuric acid signal across this period as measured by NO₃⁻ CI-API-ToF showed strong
220 midday peaks, with signal highest on 24/06/2017 and 25/06/2017. The meteorological data are
221 shown in Figure S2 alongside condensation sink (CS). The conditions were generally warm and
222 humid, with temperature reaching its maximum on 25/06/2017, with a peak hourly temperature of
223 31°C. High temperatures were seen on 21/06 and 24/06 also, of 30°C and 26°C respectively.

224

225 **3.2. Gas Phase HOM Chemistry**

226 **3.2.1. Bulk chemical properties**

227 For the peaks that have had chemical formulae assigned, oxidation state of carbon, or OS_c , can be
228 used to describe their bulk oxidation chemistry. OS_c is defined as (Kroll et al., 2011)

229

$$OSc = (2 \times O:C) - H:C \quad (2)$$

230
231
232
233
234
235
236
237
238
239
240
241
242
243
244
245
246
247
248
249
250
251

This does not account for the presence of nitrate ester groups, which has been accounted for previously by subtracting five times the N:C ratio (Massoli et al., 2018), under the assumption that all nitrogen containing functionality is in the form of nitrate ester (RONO₂) groups. In Beijing, multiple sources of nitrate-containing organic compounds are seen, in the forms of amines, nitriles and heterocycles. The variation of oxidation state with carbon number (C_n) without correction for nitrate esters is plotted in Figure 1. The average oxidation state of carbon in this dataset tends to decrease with an increase to C_n , highest where $C_n = 5$, attributable both to high O:C and peak area for the peak assigned to C₅H₁₀N₂O₈ at m/Q 288. $C_n = 5$ also shows the greatest distribution of oxidation states, likely due to the high ambient concentration of isoprene and therefore its many oxidation products being of high enough signal for many well resolved peaks to be seen in this dataset. It is worth noting that some of the ions plotted here may not form through peroxy radical autoxidation, such as C₅H₁₀N₂O₈, which may be a second-generation oxidation product of isoprene under high NO_x (Lee et al., 2016). $C_n = 10$ and 15 also see a small increase to average oxidation number compared to their neighbours. The lower oxidation state of the larger products is likely a function of two things. First and foremost, any autoxidation mechanism must undergo more steps in order for a larger molecule to reach an equivalent O:C ratio with a smaller one, and the equivalent O:C ratio is ultimately less likely to be reached before the radical is terminated (Massoli et al., 2018). Secondly, the lower vapour pressures of these larger products will lead to their partitioning into the condensed phase more readily than the smaller, thus they are more rapidly lost (Mutzel et al., 2015).

252
253
254
255

The degrees of OSc observed here are similar to those seen in other environments such as during the SOAS campaign in 2013 in southern United States, characterised by low NO/NO₂ and high temperatures, where campaign averages of 0.3 ppb, 0.4-0.5 ppb, and 25°C respectively were

256 measured, although an additional parameter to account for nitrogen containing VOCs is included in
257 the calculation (Massoli et al., 2018). The OS_c observed in Beijing is also higher than that seen in
258 the boreal forest environment of Hyytiälä, despite extremely low NO_x concentrations, likely due to
259 low temperature conditions dominating in those conditions (Schobesberger et al., 2013). These
260 relatively similar degrees of oxidation to those seen in other, cleaner, environments are likely due to
261 an interplay between the higher temperatures facilitating rapid hydrogen abstractions (Crouse et
262 al., 2013; Quéléver et al., 2019) and the higher concentrations of NO_x , HO_2 , and other RO_2
263 molecules terminating the autoxidation sequence more efficiently (Praske et al., 2018, Rissanen,
264 2018, Garmash et al., 2019).

265
266 A mass defect plot is shown in Figure 2, which shows nominal mass plotted against mass defect for
267 all peaks in this dataset. Mass defect is defined as the ion mass minus integer mass. This is shown
268 for two separate daytime periods, one where nucleation was not occurring and HOM concentrations
269 are lower (10:30 – 12:00 23/06/2017) and one where nucleation was occurring under high HOM
270 concentrations (10:30 – 12:00 25/06/2017). The band of lower mass defect is characterised by a
271 number of large peaks with high signal, for example, at m/Q 436 the ion $(C_2H_7N)_2(H_2SO_4)_2HSO_4^-$.
272 The upper component of the mass defect is dominated by organic compounds, the upper end of
273 more positive mass defect is occupied by molecules with more 1H (mass defect 7.825 mDa) and ^{14}N
274 (mass defect 3.074 mDa). The end of less positive mass defect has lower 1H and more ^{16}O (mass
275 defect -5.085 mDa); alternatively put, the mass defect reflects the variation in OS_c . The organic
276 components with more positive mass defects will be more volatile than their lower mass defect
277 counterparts as they will contain fewer oxygen functionalities (Tröstl et al., 2016, Stolzenburg et al.,
278 2018). These higher volatility products may still contribute to larger size particle growth. The more
279 negative mass defect components will be those of greater O:C and therefore lower volatility,
280 LVOCs, and the yet larger and more oxidised components, ELVOCs (Tröstl et al., 2016). During
281 the nucleation period, the signal intensity for the species in the upper band of more negative mass

282 defect have the most marked increase in concentration, with significantly less difference >500 m/Q.
283 This region 200-400 m/Q will contain most of the >C₅ monomer HOMs seen in this dataset.

284

285 **3.2.2. Diurnal trends of HOMs**

286 Temporal trends of HOMs in the urban atmosphere can reveal their sources and behaviour in the
287 atmosphere. Most of the HOM species peak in the daytime. These species all follow a similar
288 diurnal trend, as shown in Figure 3. Both the concentrations of O₃ and OH· are high during the
289 summer period in Beijing (although the nitrate chemical ionisation technique is not sensitive to all
290 OH· oxidation products (Berndt et al., 2015)). Figure S1 shows the time series of concentrations of
291 NO which is considered a dominant peroxy radical terminator of particular importance in the
292 polluted urban environment (Khan et al., 2015). Radicals such as HO₂· and RO₂· also typically peak
293 during daytime. The HOM components peaking in the daytime are presumed to be the oxidation
294 products of a mixture of anthropogenic and biogenic components, such as alkylbenzenes,
295 monoterpenes and isoprene. The oxidation of monoterpenes, specifically the monoterpene α -pinene,
296 has been the subject of extensive study recently, with the O₃-initiated autoxidation sequence being
297 the best characterised (Ehn et al., 2014; Jokinen et al., 2014; Kurtén et al., 2015; Kirkby et al.,
298 2016); ozonolysis of α -pinene opens the ring structure and produces a RO₂· radical (Kirkby et al.,
299 2016). In the case of aromatics, OH· addition to the ring and the subsequently formed bicyclic
300 peroxy radical is the basis for the autooxidation of compounds such as xylenes and
301 trimethylbenzenes (Molteni et al., 2018; Wu et al., 2017).

302

303 The identified compounds have been roughly separated into several categories, each of these plotted
304 in Figure 3. Figure 3 shows the separation of components into non-nitrogen containing HOMs, and
305 nitrogen containing HOMs, or organonitrates (ONs). The ON signal is much higher than that of the
306 HOM, attributable in part to a few ions of high signal, such as the isoprene organonitrate
307 C₅H₁₀N₂O₈. A few similar structural formulae are seen (C₅H₁₀N₂O₆, C₅H₁₁NO₆, C₅H₁₁NO₇, etc),

308 some of which have been identified as important gas phase oxidation products of isoprene under
309 high NO_x conditions (Xiong et al., 2015), and their contribution to SOA has been explored
310 previously (Lee et al., 2016). A high nitrophenol signal is also seen, $\text{C}_6\text{H}_5\text{NO}_3$. The signal for HOM
311 compounds is less dominated by a few large ions. The prevalence of ON compounds points towards
312 the important role of NO_x as a peroxy radical terminator, with the probability for the $\text{RO}_2 + \text{NO}_x$
313 reaction to produce nitrate ester compounds increasing with the size of the RO_2 molecule (Atkinson
314 et al., 1982). The NO_x concentrations in urban Beijing are approximately a factor of 10 higher than
315 seen at the Hyytiälä station in Finland as reported by Yan et al. (2016), and hence it is expected to
316 be a more significant peroxy radical terminator.

317

318 Despite the very large fluxes of anthropogenic organic pollutants in Beijing, biogenic emissions are
319 still an important source of reactive VOCs in the city, with abundant isoprene oxidation products
320 observed (see above), as well as monoterpene monomers ($\text{C}_{10}\text{H}_{16}\text{O}_9$, $\text{C}_{10}\text{H}_{15}\text{O}_9\text{N}$) and some dimer
321 products ($\text{C}_{20}\text{H}_{30}\text{O}_{11}$, $\text{C}_{20}\text{H}_{31}\text{O}_{11}\text{N}$). The time series of the signals of all C_5 , C_{10} and C_{20} molecules is
322 plotted in Figure 3b, with C_5 species assumed to be isoprene dominated, C_{10} and C_{20} assumed to be
323 monoterpene dominated. Signals for isoprene oxidation products are higher, with abundant
324 isoprene nitrate and dinitrate products. C_{10} products show similar behaviour, with, for example,
325 several $\text{C}_{10}\text{H}_{15}\text{O}_x\text{N}$ $x = 5-9$ compounds seen. The C_{20} signal intensities are low, and follow the
326 general formula $\text{C}_{20}\text{H}_x\text{O}_y\text{N}_z$, where $x = 26-32$, $y = 7-11$ and $z = 0-2$; in Figure 3 the signal for C_{20}
327 compounds has been multiplied by a factor of 50 for visibility. The low signals reflect the lack of
328 RO_2 cross reactions necessary for the production of these accretion products.

329

330 Other identified peaks are plotted in Figure 3c. The C_2 - C_4 components are summed together, these
331 being small organic acids such as malonic acid and oxalic acid, as well as products such as
332 $\text{C}_4\text{H}_7\text{O}_6\text{N}$. Malonic acid is the most prominent here, seen both as an NO_3^- adduct ($\text{C}_3\text{H}_4\text{O}_4\text{NO}_3^-$) and
333 a proton transfer product ($\text{C}_3\text{H}_3\text{O}_4^-$) at a ratio of around 2:3. Measurements of particle phase

334 dicarboxylic acids in cities typically show greater concentrations of of oxalic acid than malonic (Ho
335 et al., 2016), and these acids are primarily produced in the aqueous phase (Bikkina et al., 2014).
336 Primary sources of dicarboxylic acid include fossil fuel combustion (Kawamura and Kaplan, 1987)
337 and biomass burning (Narukawa et al., 1999), which are both plentiful in urban Beijing. The C₆-C₉
338 components are assumed to be dominated by oxidation products of alkylbenzenes such as C₈H₁₂O₅,
339 although fragments of other compounds, i.e., monoterpenes, can also occupy this region (Isaacman-
340 Vanwertz et al., 2018). It is assumed the majority of the signal for these peaks come from
341 alkylbenzenes. This assumption is supported by the relative signal intensity ratios of the oxygen
342 numbers of monomer C₈H₁₂O_n compounds being similar to those seen for xylene oxidation products
343 in previous work (Molteni et al., 2018). The largest fraction, C₁₁ through C₁₈, includes the larger
344 compounds, oxidation products of larger aromatics, or products of the cross reaction of smaller RO₂
345 radicals. Here they are grouped without more sophisticated disaggregation as they all follow much
346 the same time series, species such as C₁₁H₁₁O₈N following the same temporal trends as C₁₅H₁₆O₉
347 and C₁₆H₂₄O₁₂.

348

349 Nearly all ions with the exception of the larger compounds attributed to the cross reaction of C₁₀
350 monomers follow similar temporal patterns, with the majority of peaks occurring in the daytime.
351 This reflects the importance of the concentration of atmospheric oxidants. Some selected oxidation
352 products are plotted against their precursor VOCs in Figure 4. The concentration of isoprene is
353 plotted against the signal of a nitrate HOM product, C₅H₉NO₆ (Xiong et al., 2015; Lee et al., 2016),
354 while monoterpenes are plotted against C₁₀H₁₆O₉ (Ehn et al., 2014; Berndt et al., 2016; Yan et al.,
355 2016; Kirkby et al., 2016; Massoli et al., 2018), and C₂-benzenes against C₈H₁₂O₆ (Molteni et al.,
356 2018; Wang et al., 2017). The first half of the time series shows little correlation between the VOC
357 species and the resultant oxidation products, while isoprene, monoterpenes and C₂-benzenes follow
358 their usual diurnal cycles, isoprene having the most distinct with a strong midday peak. The latter
359 two days, however, show similar and coinciding peaks in both the VOCs and HOMs - HOMs show

360 afternoon peaks on both days, and an initial shelf on the final half day. The $C_5H_9NO_6$ peak follows
361 some of the peaks of the isoprene, but not all (e.g., morning shelf of isoprene on 24/06).
362 Concentrations of isoprene do not seem to determine directly the signal of HOM, as the day with
363 the lowest isoprene of all is the day with highest $C_5H_9NO_6$. The $C_{10}H_{16}O_9$ trace has coincidental
364 peaks with the monoterpene trace also, including two 4-hour separated simultaneous peaks on
365 25/06. The peaks in the concentrations of C_2 -benzenes are nearly synchronous with the peaks in
366 $C_8H_{12}O_6$, for which the data exhibit a strong mid afternoon peak likely due to the lack of an efficient
367 ozonolysis reaction pathway; the main oxidant of C_2 -benzenes is the OH radical. Trends of both C_3
368 benzenes and their HOMs are much the same as C_2 benzenes as discussed above, pointing to similar
369 sources and oxidation chemistries. . The concentration of precursor VOC is likely a driving force in
370 the identity and quantity of various HOM products, but not the sole determinant, as while there are
371 simultaneous peaks of VOCs and HOMs, both the condensation sink and oxidant concentrations
372 also influence HOM product signals.

373
374 The first half of campaign measurements is marked by an episode of low HOM signals. A diurnal
375 cycle still exists but it is weak. The radiation intensity was significantly lower on these prior days
376 than it was on the 24th. No data is available for the final period of measurement. Ozone is higher on
377 the prior measurement days with lower HOM signals (see Figure S1). Little agreement is seen
378 between VOC concentration and HOM signals on these days. The condensational sinks are roughly
379 similar to those on days of higher HOM concentrations, but temperature and solar radiation are
380 much lower. HOM formation is largely dependent upon VOC concentration, oxidant concentration
381 (which will be lower if solar radiation is lower, especially in the case of OH, the main oxidant of
382 aromatic species especially), and temperature (as H-shifts are highly temperature dependent)
383 (Quéléver et al., 2019), as well as losses by RO_2 termination before a molecule can become HOM,
384 and losses to condensational sink. The low HOM concentration is likely due to these lower
385 temperatures, and weaker solar radiation not facilitating HOM formation.

386 The C₂₀ compounds plotted in Figure 3b show no strong diurnal sequence, contrasting with other
387 HOMs. We can presume that all C₂₀ compounds identified are the result of the reaction of two
388 monoterpene C₁₀ RO₂ radicals, a reasonable assumption as all identified C₂₀ species follow the
389 general formula outlined for these reactions (C₂₀H₂₈₋₃₂O₆₋₁₆). The formation of C₂₀ dimers is
390 dependent upon two processes, initial oxidation of monoterpenes, and RO₂-RO₂ termination. Initial
391 oxidation is contingent upon oxidant concentration, which is highest in the daytime, and RO₂-RO₂
392 termination is contingent upon the probability of the molecular collision between the RO₂
393 molecules occurring before other radical termination (i.e., RO₂-NO_x, or RO₂-HO₂). There is likely
394 a strong diurnal sequence in the dominant RO₂ termination mechanisms across the day period, and
395 the combination of the two factors discussed above results in there being no strong diurnal trend in
396 these molecules. A lower oxidant concentration at night results in less RO₂ molecules, but less NO
397 and HO₂ results in a greater chance for those RO₂ molecules to dimerise (Rissanen, 2018, Garmash
398 et al., 2019). As the levels of NO_x in Beijing fall, the peroxy radical termination reactions will be
399 less probable compared to continued autoxidation (Praske et al., 2018), and it is expected that more
400 oxidised HOM products will be seen with lower volatilities and therefore a greater potential
401 contribution to earlier stage particle formation and growth.

402

403 3.3. New Particle Formation

404 Nearly all the signal intensity in the CI-API-ToF instrument arises from molecules charged by NO₃⁻,
405 therefore plotting the unit mass resolution data (the data gained by integrating over the entire area at
406 each m/Q integer) against time describes simply the evolution of oxidised organic molecules, acids
407 and their molecular clusters both with each other and stabilising amine species. This is done in
408 Figure 5. As the signal intensity varies by factors of 10 from mass to mass, each value has been
409 normalised so they have maxima at 1. This has been done separately for two days for clarity, as the
410 signal intensity also varies from day to day. PSM data for these two days is plotted in Figure 5 also,
411 with both total particle count >1.30 nm in black and the number difference between the lower and

412 upper size cuts (1.30 and 1.84 nm) in blue, which shows the number of particles between these
413 sizes. The relationship between mass and electrical mobility diameter can be defined thus (Tammet,
414 1995),

$$415 \quad d_e = \left(\frac{6m}{\pi\rho}\right)^{\frac{1}{3}} + d_g \quad (3)$$

416

417 where d_e is the electrical mobility diameter of the cluster or particle, m is the mass of the cluster or
418 particle expressed in kg, ρ is the density and d_g is the effective gas diameter, determined to be 0.3
419 nm for smaller particles (Larriba et al., 2011). We can use this to draw a comparison between the
420 PSM and CI-APi-ToF measurements. If a density of 1.2 g cm^{-3} is assumed, then once molecular
421 clusters reach the $>400 \text{ m/Q}$ range, they will be seen in the lowest size cut of the PSM, or $>700 \text{ m/Q}$
422 if a density of 2.0 g cm^{-3} is assumed. A full table of densities is provided in the Supplementary
423 Information.

424

425 A burst in the signal seen by the CI-APi-TOF occurs first in the late morning in the top panel of
426 Figure 5, and this is at the same time as peaks begin to rise in the identified HOMs (see Figure 3).
427 Here, the PSM is not available due to an instrumental fault until 16:00; however, at that point, an
428 elevation to particle count and a large elevation to cluster count can be seen. Moving into the
429 evening period, the mass contour shows peaks to larger masses $>400 \text{ m/Q}$. This is likely dimerised
430 compounds and products of NO_3^- chemistry with little contribution to newly forming particles, but
431 still sensitive to chemical ionisation by NO_3^- . Many of these peaks cannot be assigned due to
432 uncertainties in the structural formula assignment for higher mass peaks, as the number of possible
433 dimerised compounds is many, being the combination of most possible RO_2 radicals. Graphically,
434 these are over-represented in Figure 5 due to the normalisation, their signals (especially $>500 \text{ m/Q}$)
435 are much lower than the signals $<400 \text{ m/Q}$.

436

437 The second day plotted in the lower panel of Figure 5 (25/06/2017) shows a strong afternoon peak
438 to the HOMs (for most HOMs, stronger than that on the day prior). Particle formation is shown in
439 the PSM data. A strong midday peak to particle number is seen with two distinct peaks to cluster
440 count. These two peaks are not coincidental with the two peaks to HOM signal (i.e., nitrogen-
441 containing HOMs in Figure 3a peaking at 11:00 and 16:00). Sulfuric acid, however, does peak
442 synchronously with the particle number count. Sulfuric acid is plotted across the contour plot in
443 Figure 6, where PSM data is also shown in the bottom panel. The peak to CI-APi-TOF mass signal,
444 visible in Figure 5 occurs at around 12:00/13:00, peaks in the PSM cluster count occur at 10:00 and
445 13:00. Peaks in mass up to 550 m/Q are seen in the CI-APi-ToF at 13:00. Assuming the density of
446 these species is $\leq 1.6 \text{ g cm}^{-3}$ then these will be suitably sized to be grown in the PSM saturator..
447 These newly formed particles then go on to grow and contribute significantly to the larger particle
448 count (Figure S3). As initial particle formation coincides with sulfuric acid signal peaks and before
449 HOM signals peak, it can be assumed on these days, the HOM contribution to the initial particle
450 formation is modest.

451

452 There is recent strong evidence to suggest that the driving force of the earliest stages of particle
453 formation in urban Shanghai is from sulfuric acid and C₂-amines (Yao et al., 2018), and the
454 coincidental peaks of sulfuric acid with new particles as seen in Figure 6 suggest a similar
455 behaviour. Dimethylamine (DMA) can efficiently stabilise the sulfuric acid clusters (Almeida et al.,
456 2013). Here, few larger sulfuric acid-DMA clusters were visible in the dataset, as seen in the work
457 by Yao et al., 2018, although five sulfuric acid-dimethylamine (SA-DMA) ions were observed, the
458 others were likely too low in signal to be confidently resolved from their neighbouring peaks;
459 however, clusters of up to 4 sulfuric acid molecules and 3 dimethylamine molecules were seen, with
460 similar diurnal trends to sulfuric acid. The scarcity of SA-DMA clusters is likely due to
461 instrumental conditions, rather than their absence in the atmosphere. The nitrate chemical ionisation
462 system tends to evaporate amine compounds upon charging, and as specific voltage-tuning setups

463 can lend themselves towards preservation or breakage of molecular clusters, the signal for larger
464 sulfuric acid clusters was also very weak. The formation of HOM-sulfuric acid clusters is unlikely
465 under atmospheric conditions (Elm et al., 2017) and few of these were observed. Signals of HOMs
466 seem to coincide with later particle growth; it can be expected that HOM molecules make a more
467 significant contribution to particle growth than to early particle formation, with the largest and most
468 oxidised being involved in early growth, and the smaller and less oxidised contributing to later
469 growth as the necessary vapour pressure properties become less demanding.

470

471 **4. CONCLUSIONS**

472 The average degree of HOM oxidation in Beijing is comparable with that seen in other
473 environments. Rapid intramolecular hydrogen shifts during autoxidation due to the higher
474 temperatures are probably offset by the frequent termination reactions due to high NO_x
475 concentrations. OS_c values seem to be marginally higher for biogenic species.

476

477 The temporal trend of nearly every HOM shows afternoon or evening maxima. Both O_3 and $\text{OH}\cdot$
478 have high daytime concentrations and these likely drive the initial oxidation steps. The species
479 arising from alkylbenzene precursors show sharper afternoon peaks, probably since their oxidation
480 is $\text{OH}\cdot$ dominated. Many of the rest of the peaks, coming from largely BVOC precursors show
481 broader daytime peaks, being influenced by O_3 also. There seems to be no direct link between VOC
482 concentrations and HOM signals, with days of lower precursor VOC sometimes having higher
483 HOM signals and vice versa.

484

485 Initial particle formation coincides with peak sulfuric acid signals, while the growth of the particles
486 correlates more closely with the signals of HOMs. This is very similar to behaviour observed in a
487 study of NPF in Shanghai which was attributed to sulfuric acid-dimethylamine-water nucleation
488 with condensing organic species contributing to particle growth (Yao et al., 2018), and this is

489 further backed up by numerous SA-DMA clusters present in this dataset. The freshly formed
490 particles grow and contribute significantly to total particle loading. This is visible when the unit
491 mass CI-APi-ToF data is plotted as a contour plot, and further to this is visible in the PSM data,
492 with bursts to both total number count >1.30 nm and the number of molecular clusters between 1.30
493 and 1.84 nm. As NO_x levels fall in Beijing due to traffic emission control measures being enforced
494 it is likely that autoxidation will become increasingly significant in the new particle formation
495 processes. The number of molecules detected by the NO₃ CIMS is undoubtedly many more than
496 have had formulae assigned here, but to identify more requires a more sophisticated data
497 deconvolution.

498

499 **DATA ACCESSIBILITY**

500 Data supporting this publication are openly available from the UBIRA eData repository at
501 <https://doi.org/10.25500/edata.bham.00000304>

502

503 **AUTHOR CONTRIBUTIONS**

504 The study was conceived and planned by RMH and ZS. DCSB and JB set up and operated the
505 main instrumental measurements, and JB prepared the first draft of the paper and responded to
506 comments from RMH and ZS. CNH and WJA contributed the hydrocarbon data and provided
507 comments on the draft manuscript, and ES and JL contributed the gas phase pollutant data.

508

509 **COMPETING INTERESTS**

510 The authors have no conflict of interests.

511

512 **ACKNOWLEDGMENTS**

513 This work was part of the APHH-Beijing programme funded by the UK Natural Environment
514 Research Council (NE/N007190/1) and the Natural Sciences Funding Council of China. It was

515 additionally facilitated by the National Centre for Atmospheric Science ODA national capability
516 programme ACREW (NE/R000034/1), which is supported by NERC and the GCRF. We thank
517 Professor X.M Wang from the Guangzhou Institute of Geochemistry, Chinese Academy of
518 Sciences, Brian Davison from Lancaster University and Ben Langford, Eiko Nemitz, Neil
519 Mullinger and other staff from the Centre for Ecology and Hydrology, Edinburgh for assistance
520 with the VOC measurements and associated infrastructure.

521

522 **REFERENCES**

523

524 Alam, A., Shi, J.P., Harrison R.M.: Observations of new particle formation in urban air, *J. Geophys.*
525 *Res.*, 108, 4093-4107, doi:10.1029/2001JD001417, 2003

526

527 Almeida, J., Schobesberger, S., Kürten, A., Ortega, I. K., Kupiainen-Määttä, O., Praplan, A. P.,
528 Adamov, A., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Donahue, N.
529 M., Downard, A., Dunne, E., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Guida, R.,
530 Hakala, J., Hansel, A., Heinritzi, M., Henschel, H., Jokinen, T., Junninen, H., Kajos, M.,
531 Kangasluoma, J., Keskinen, H., Kupc, A., Kurtén, T., Kvashin, A. N., Laaksonen, A., Lehtipalo, K.,
532 Leiminger, M., Leppä, J., Loukonen, V., Makhmutov, V., Mathot, S., McGrath, M. J., Nieminen,
533 T., Olenius, T., Onnela, A., Petäjä, T., Riccobono, F., Riipinen, I., Rissanen, M., Rondo, L.,
534 Ruuskanen, T., Santos, F. D., Sarnela, N., Schallhart, S., Schnitzhofer, R., Seinfeld, J. H., Simon,
535 M., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A., Tröstl, J., Tsagkogeorgas, G., Vaattovaara,
536 P., Viisanen, Y., Virtanen, A., Vrtala, A., Wagner, P. E., Weingartner, E., Wex, H., Williamson, C.,
537 Wimmer, D., Ye, P., Yli-Juuti, T., Carslaw, K. S., Kulmala, M., Curtius, J., Baltensperger, U.,
538 Worsnop, D. R., Vehkamäki, H., and Kirkby, J.: Molecular understanding of sulphuric acid-amine
539 particle nucleation in the atmosphere, *Nature*, 502, 359-363, 2013.

540

541 Atkinson, R., Aschmann, S. M., Carter, W. P. L., Winer, A. M., and Pitts, J. N.: Alkyl nitrate
542 formation from the nitrogen oxide (NO_x)air photooxidations of C2-C8 n-alkanes, *J. Phys. Chem.*,
543 86, 4563-4569, 1982.

544

545 Berndt, T., Richters, S., Kaethner, R., Voigtländer, J., Stratmann, F., Sipilä, M., Kulmala, M., and
546 Herrmann, H.: Gas-Phase Ozonolysis of Cycloalkenes: Formation of Highly Oxidized RO₂
547 Radicals and Their Reactions with NO, NO₂, SO₂, and Other RO₂ Radicals, *J. Phys. Chem. A.*,
548 119, 10336-10348, 2015.

549

550 Bikkina, S., Kawamura, K., Miyazaki, Y., and Fu, P.: High abundances of oxalic, azelaic, and
551 glyoxylic acids and methylglyoxal in the open ocean with high biological activity: Implication for
552 secondary OA formation from isoprene, *Geophys. Res. Lett.*, 41, 3649–3657,
553 doi:10.1002/2014GL059913, 2014.

554 Berndt, T., Richters, S., Jokinen, T., Hyttinen, N., Kurtén, T., Otkjær, R. V., Kjaergaard, H. G.,
555 Stratmann, F., Herrmann, H., Sipilä, M., Kulmala, M., and Ehn, M.: Hydroxyl radical-induced
556 formation of highly oxidized organic compounds, *Nature Comm.*, 7, 20
557 <https://doi.org/10.1038/ncomms13677>, 2016.

558

559 Bianchi, F., Garmash, O., He, X., Yan, C., Iyer, S., Rosendahl, I., Xu, Z., Rissanen, M. P., Riva, M.,
560 Taipale, R., Sarnela, N., Petäjä, T., Worsnop, D. R., Kulmala, M., Ehn, M., and Junninen, H.: The
561 role of highly oxygenated molecules (HOMs) in determining the composition of ambient ions in the
562 boreal forest, *Atmos. Chem. Phys.*, 17, 13819-13831, 2017.

563

564 Brook, R. D., Rajagopalan, S., Pope, C. A., Brook, J. R., Bhatnagar, A., Diez-Roux, A. V., Holguin,
565 F., Hong, Y., Luepker, R. V., Mittleman, M. A., Peters, A., Siscovick, D., Smith, S. C., Whitsel, L.,
566 and Kaufman, J. D.: Particulate matter air pollution and cardiovascular disease: An update to the
567 scientific statement from the american heart association, *Circulation*, 121, 2331-2378, 2010.

568

569 Chu, B., Kerminen, V.-M., Bianchi, F., Yan, C., Petaja, T., and Kulmala, M.: Atmospheric new
570 particle formation in China, *Atmos. Chem. Phys.*, 19, 115-138, 2019.

571

572 Crouse, J. D., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., and Wennberg, P. O.: Autoxidation
573 of organic compounds in the atmosphere, *J. Phys. Chem. Lett.*, 4, 3513-3520, 2013.

574
575 Cubison, M. J. and Jimenez, J. L.: Statistical precision of the intensities retrieved from constrained
576 fitting of overlapping peaks in high resolution mass spectra, *Atmos. Meas. Tech.*, 8, 2333-2345,
577 2015.
578
579 Delfino, R. J., Sioutas, C., and Malik, S.: Potential role of ultrafine particles in associations between
580 airborne particle mass and cardiovascular health, *Environ. Health Perspect.*, 113, 934-946, 2005.
581
582 Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F.,
583 Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., Jokinen, T.,
584 Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T., Nielsen,
585 L. B., Jørgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T., Petäjä, T.,
586 Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large
587 source of low-volatility secondary organic aerosol, *Nature*, 506, 476-479, 2014.
588
589 Elm, J., Myllys, N., and Kurtén, T.: What is Required for Highly Oxidized Molecules to Form
590 Clusters with Sulfuric Acid?, *J. Phys. Chem. A*, 121, 4578-4587, 2017.
591
592 Garmash, O., Rissanen, M. P., Pullinen, I., Schmitt, S., Kausiala, O., Tillmann, R., Percival, C.,
593 Bannan, T. J., Priestley, M., Hallquist, Å. M., Kleist, E., Kiendler-Scharr, A., Hallquist, M., Berndt,
594 T., McFiggans, G., Wildt, J., Mentel, T., and Ehn, M.: Multi-generation OH oxidation as a source
595 for highly oxygenated organic molecules from aromatics, *Atmos. Chem. Phys. Discuss.*,
596 <https://doi.org/10.5194/acp-2019-582>, in review, 2019.
597
598 Gong, Y., Hu, M., Cheng, Y., Su, H., Yue, D., Liu, F., Wiedensohler, A., Wang, Z., Kalesse, H.,
599 Liu, S., Wu, Z., Xiao, K., Mi, P., and Zhang, Y.: Competition of coagulation sink and source rate:
600 New particle formation in the Pearl River Delta of China, *Atmos. Environ.*, 44, 3278-3285, 2010.
601
602 Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M., Zeng, L.,
603 Molina, M. J., and Zhang, R.: Elucidating severe urban haze formation in China., *PNAS*, 111,
604 17373- 17378, 2014.
605
606 Ho, K. F., Lee, S. C., Ho, S. S. H., Kawamura, K., Tachibana, E., Cheng, Y. and Zhu, T.: Dicarboxylic
607 acids, ketocarboxylic acids, α -dicarbonyls, fatty acids, and benzoic acid in urban aerosols collected
608 during the 2006 Campaign of Air Quality Research in Beijing (CAREBeijing-2006), *J. Geophys. Res.*
609 *Atmos.*, 115(19), 1–14, doi:10.1029/2009JD013304, 2010.
610
611 Hyttinen, N., Kupiainen-Määttä, O., Rissanen, M. P., Muuronen, M., Ehn, M., and Kurtén, T.:
612 Modeling the Charging of Highly Oxidized Cyclohexene Ozonolysis Products Using Nitrate-Based
613 Chemical Ionization, *J. Phys. Chem.*, A, 119, 6339-6345, 2015.
614
615 Isaacman-Vanwertz, G., Massoli, P., O'Brien, R., Lim, C., Franklin, J. P., Moss, J. A., Hunter, J. F.,
616 Nowak, J. B., Canagaratna, M. R., Misztal, P. K., Arata, C., Roscioli, J. R., Herndon, S. T., Onasch,
617 T. B., Lambe, A. T., Jayne, J. T., Su, L., Knopf, D. A., Goldstein, A. H., Worsnop, D. R., and Kroll,
618 J. H.: Chemical evolution of atmospheric organic carbon over multiple generations of oxidation,
619 *Nature Chem.*, 10, 462-468, <https://doi.org/10.1038/s41557-018-0002-2>, 2018
620
621
622
623 Jokinen, T., Sipilä, M., Junninen, H., Ehn, M., Lönn, G., Hakala, J., Petäjä, T., Mauldin, R. L.,
624 Kulmala, M., and Worsnop, D. R.: Atmospheric sulphuric acid and neutral cluster measurements

625 using CI-API-TOF, *Atmos. Chem. Phys.*, 12, 4117-4125, 2012.

626

627 Jokinen, T., Sipilä, M., Richters, S., Kerminen, V. M., Paasonen, P., Stratmann, F., Worsnop, D.,
628 Kulmala, M., Ehn, M., Herrmann, H., and Berndt, T.: Rapid autoxidation forms highly oxidized
629 RO₂ radicals in the atmosphere, *Angewandte Chemie - International Edition*, 53, 14596-14600,
630 <https://doi.org/10.1002/anie.201408566>, 2014.

631

632 Junninen, H., Ehn, M., Petäjä, Luosujärvi, L., Kotiaho, T., Kostianen, R., Rohner, U., Gonin, M.,
633 Fuhrer, K., Kulmala, M., and Worsnop, D. R.: A high-resolution mass spectrometer to measure
634 atmospheric ion composition, *Atmos. Meas. Tech.*, 3, 1039-1053, 2010.

635

636 Kerminen, V. M., Paramonov, M., Anttila, T., Riipinen, I., Fountoukis, C., Korhonen, H., Asmi, E.,
637 Laakso, L., Lihavainen, H., Swietlicki, E., Svenningsson, B., Asmi, A., Pandis, S. N., Kulmala, M.,
638 and Petäjä, T.: Cloud condensation nuclei production associated with atmospheric nucleation: A
639 synthesis based on existing literature and new results, *Atmos. Chem. Phys.*, 12, 12037-12059, 2012.

640

641 Kawamura, K. and Kaplan, I. R.: Motor Exhaust Emissions as a Primary Source for Dicarboxylic
642 Acids in Los Angeles Ambient Air, *Environ. Sci. Technol.*, 21(1), 105–110,
643 [doi:10.1021/es00155a014](https://doi.org/10.1021/es00155a014), 1987.

644

645 Khan, M., Cooke M, Utembe, S., Archibald A., Derwent, R., Jenkin, M., Morris, W., South, N.,
646 Hansen, J., Francisco, J., Percival, C., Shallcross, D.: Global analysis of peroxy radicals and peroxy
647 radical-water complexation using the STOCHEM-CRI global chemistry and transport model,
648 *Atmos. Environ.*, 106, 278-287, 2015.

649

650 Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné, S.,
651 Ickes, L., Kürten, A., Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S.,
652 Tsagkogeorgas, G., Wimmer, D., Amorim, A., Bianchi, F., Breitenlechner, M., David, A.,
653 Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S., Hansel, A., Hauser, D., Jud, W.,
654 Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R.,
655 Makhmutov, V., Mathot, S., Mikkilä, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A.,
656 Pereira, P., Petäjä, T., Schnitzhofer, R., Seinfeld, J. H., Sipilä, M., Stozhkov, Y., Stratmann, F.,
657 Tomé, A., Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E.,
658 Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger, U., and Kulmala, M.:
659 Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol nucleation,
660 *Nature*, 476, 429-435, <https://doi.org/10.1038/nature10343>, 2011.

661

662 Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon,
663 M., Yan, C., Almeida, J., Trostl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim,
664 A., Bernhammer, A. K., Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X., Craven, J., Dias, A.,
665 Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T.,
666 Junninen, H., Kangasluoma, J., Kim, J., Krapf, M., Kurten, A., Laaksonen, A., Lehtipalo, K.,
667 Makhmutov, V., Mathot, S., Molteni, U., Onnela, A., Perakyla, O., Piel, F., Petaja, T., Praplan, A.
668 P., Pringle, K., Rap, A., Richards, N. A., Riipinen, I., Rissanen, M. P., Rondo, L., Sarnela, N.,
669 Schobesberger, S., Scott, C. E., Seinfeld, J. H., Sipilä, M., Steiner, G., Stozhkov, Y., Stratmann, F.,
670 Tomé, A., Virtanen, A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E., Wimmer, D.,
671 Winkler, P. M., Ye, P., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R.,
672 Baltensperger, U., Kulmala, M., Carslaw, K. S., and Curtius, J.: Ion-induced nucleation of pure
673 biogenic particles, *Nature*, 533, 521-526, <https://doi.org/10.1038/nature17953>, 2016.

674

675 Kroll, J. H., Donahue, N. M., Jimenez, J. L., Kessler, S. H., Canagaratna, M. R., Wilson, K. R.,
676 Altieri, K. E., Mazzoleni, L. R., Wozniak, A. S., Bluhm, H., Mysak, E. R., Smith, J. D., Kolb, C. E.,

677 and Worsnop, D. R.: Carbon oxidation state as a metric for describing the chemistry of atmospheric
678 organic aerosol, *Nature Chemistry*, 3, 133-139, <https://doi.org/10.1038/nchem.948>, 2011.

679

680 Kulmala, M., Petäjä, T., Mönkkönen, P., Koponen, I. K., Dal Maso, M., Aalto, P. P., Lehtinen, K.
681 E. J., and Kerminen, V.-M.: On the growth of nucleation mode particles: source rates of
682 condensable vapor in polluted and clean environments, *Atmos. Chem. Phys.*, 5, 409-416, 2005.

683

684 Kulmala, M., Petäjä, T., Nieminen, T., Sipilä, M., Manninen, H. E., Lehtipalo, K., Dal Maso, M.,
685 Aalto, P. P., Junninen, H., Paasonen, P., Riipinen, I., Lehtinen, K. E. J., Laaksonen, A., and
686 Kerminen, V.-M.: Measurement of the nucleation of atmospheric aerosol particles, *Nature*
687 *Protocols*, 7, 1651-1667, <https://doi.org/10.1038/nprot.2012.091>, 2012.

688 Kürten, A., Rondo, L., Ehrhart, S., and Curtius, J.: Calibration of a chemical ionization mass
689 spectrometer for the measurement of gaseous sulfuric acid, *J. Phys. Chem., A*, 116, 6375-6386.

690

691 Kurtén, T., Rissanen, M. P., Mackeprang, K., Thornton, J. A., Hyttinen, N., Jørgensen, S., Ehn, M.,
692 and Kjaergaard, H. G.: Computational Study of Hydrogen Shifts and Ring-Opening Mechanisms in
693 α -Pinene Ozonolysis Products, *J. Phys. Chem., A*, 119, 11366-11375, 2015.

694

695 Kurtén, T., Tiusanen, K., Roldin, P., Rissanen, M., Luy, J. N., Boy, M., Ehn, M., and Donahue, N.:
696 α -Pinene autoxidation products may not have extremely low saturation vapor pressures despite high
697 O:C ratios, *J. Phys. Chem., A*, 120, 2569-2582, 2016.

698

699 Larriba, C., Hogan, C. J., Attoui, M., Borrajo, R., Garcia, J. F., and De La Mora, J. F.: The
700 mobility-volume relationship below 3.0 nm examined by tandem mobility-mass measurement,
701 *Aerosol Sci. Technol.*, 45, 453-467, 2011.

702

703 Lee, B. H., Mohr, C., Lopez-Hilfiker, F. D., Lutz, A., Hallquist, M., Lee, L., Romer, P., Cohen, R.
704 C., Iyer, S., Kurtén, T., Hu, W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Xu, L., Ng, N. L.,
705 Guo, H., Weber, R. J., Wild, R. J., Brown, S. S., Koss, A., de Gouw, J., Olson, K., Goldstein, A. H.,
706 Seco, R., Kim, S., McAvey, K., Shepson, P. B., Starn, T., Baumann, K., Edgerton, E. S., Liu, J.,
707 Shilling, J. E., Miller, D. O., Brune, W., Schobesberger, S., D'Ambro, E. L., and Thornton, J. A.:
708 Highly functionalized organic nitrates in the southeast United States: Contribution to secondary
709 organic aerosol and reactive nitrogen budgets, *PNAS*, 113, 1516-1521, 2016.

710

711 Lehtipalo, K., Yan, C., Dada, L., Bianchi, F., Xiao, M., Wagner, R., Stolzenburg, D., Ahonen, L. R.,
712 Amorim, A., Baccarini, A., Bauer, P. S., Baumgartner, B., Bergen, A., Bernhammer, A.,
713 K., Breitenlechner, M., Brilke, S., Buchholz, A., Mazon, S. B., Chen, D., Chen, X., Dias, A.,
714 Dommen, J., Draper, D. C., Duplissy, J., Ehn, M., Finkenzeller, H., Fischer, L., Frege, C., Fuchs,
715 C., Garmash, O., Gordon, H., Hakala, J., He, X., Heikkinen, L., Heinritzi, M., Helm, J. C., Hofbauer,
716 V., Hoyle, C. R., Joki-nen, T., Kangasluoma, J., Kerminen, V.-M., Kim, C., Kirkby, J., Kontkanen,
717 J., Kürten, A., Lawler, M. J., Mai, H., Mathot, S., Mauldin, R. L., Molteni, U., Nichman, L., Nie,
718 W., Niemi-nen, T., Ojdanic, A., Onnela, A., Passananti, M., Petäjä, T., Piel, S., Pospisilova, V.,
719 Quéléver, L. L. J., Rissanen, M. P., Rose, C., Sarnela, N., Schallhart, S., Schuchmann, S., Sengupta,
720 K., Simon, M., Sipilä, M., Tauber, C., Tomé, A., Tröstl, J., Väisä-nen, O., Vogel, A. L., Volkamer,
721 R., Wagner, A. C., Wang, M., Weitz, L., Wimmer, D., Ye, P., Ylisirniö, A., Zha, Q., Carslaw, K. S.,
722 Curtius, J., Donahue, N. M., Flagan, R. C., Hansel, A., Riipinen, I., Virtanen, A., Winkler, P. M.,
723 Baltensperger, U., Kulmala, M., and Worsnop, D. R.: Multicomponent new particle formation from
724 sulfuric acid, ammonia, and biogenic vapors, *Science Advances*, 4,
725 <https://doi.org/10.1126/sciadv.aau5363>, 2018.

726

727 Massoli, P., Stark, H., Canagaratna, M. R., Krechmer, J. E., Xu, L., Ng, N. L., Mauldin, R. L., Yan,
728 C., Kimmel, J., Misztal, P. K., Jimenez, J. L., Jayne, J. T., and Worsnop, D. R.: Ambient

729 Measurements of Highly Oxidized Gas-Phase Molecules during the Southern Oxidant and Aerosol
730 Study (SOAS) 2013, *ACS Earth Space Chem.*, 2, 653-672, 2018.

731

732 McMurry, P. H., Shan Woo, K., Weber, R., Chen, D.-R., and Pui, D. Y. H.: Size distributions of 3-
733 10 nm atmospheric particles: implications for nucleation mechanisms, *Philosophical Transactions*
734 *of the Royal Society A: Math., Phys. Eng. Sci.*, 358, 2625-2642, 2000.

735

736 Miller, M. R., Raftis, J. B., Langrish, J. P., McLean, S. G., Samuotrta, P., Connell, S. P., Wilson, S.,
737 Vesey, A. T., Fokkens, P. H., Boere, A. J. F., Krystek, P., Campbell, C. J., Hadoke, P. W.,
738 Donaldson, K., Cassee, F. R., Newby, D. E., Duffin, R., and Mills, N. L.: Inhaled nanoparticles
739 accumulate at sites of vascular disease, *ACS Nano*, 11, 4542-4552, 2017.

740 Møller, K. H., Tram, C. M., and Kjaergaard, H. G.: Side-by-Side Comparison of Hydroperoxide
741 and Corresponding Alcohol as HydrogenBond Donors, *J. Phys. Chem. A*, 121, 2951-2959, 2017.

742

743 Molteni, U., Bianchi, F., Klein, F., Haddad, I. E., Frege, C., Rossi, M. J., Dommen, J., and
744 Baltensperger, U.: Formation of highly oxygenated organic molecules from aromatic compounds,
745 *Atmos. Chem. Phys.*, 18, 1909-1921, 2018.

746

747 Mutzel, A., Poulain, L., Berndt, T., Iinuma, Y., Rodigast, M., Böge, O., Richters, S., Spindler, G.,
748 Sipila, M., Jokinen, T., Kulmala, M., Herrmann, H.: Highly oxidized multifunctional organic
749 compounds observed in tropospheric particles: A field and laboratory study, *Environ. Sci. Techn.*,
750 49, 7754-7761, 2015.

751

752 Myllys, N., Olenius, T., Kurtén, T., Vehkamäki, H., Riipinen, I., and Elm, J.: Effect of Bisulfate,
753 Ammonia, and Ammonium on the Clustering of Organic Acids and Sulfuric Acid, *J. Phys. Chem.*
754 *A*, 121, 4812-4824, 2017.

755

756 Narukawa, M., Kawamura, K., Takeuchi, N. and Nakajima, T.: Distribution of dicarboxylic acids
757 and carbon isotopic ratios in aerosols from 1997 Indonesian forest fires, *Geophys. Res. Lett.*,
758 26(10), 3101-3104, 1999.

759

760 Penner, J. E., Xu, L., and Wang, M.: Satellite methods underestimate indirect climate forcing by
761 aerosols., *PNAS*, 108, 13404-13408, 2011.

762

763 Praske, E., Otkjær, R. V., Crouse, J. D., Hethcox, J. C., Stoltz, B. M., Kjaergaard, H. G., and
764 Wennberg, P. O.: Atmospheric autoxidation is increasingly important in urban and suburban North
765 America, *PNAS*, 115, 64-69, 2018.

766

767 Qi, X., Ding, A., Roldin, P., Xu, Z., Zhou, P., Sarnela, N., Nie, W., Huang, X., Rusanen, A., Ehn,
768 M., Rissanen, M. P., Petäjä, T., Kulmala, M., and Boy, M.: Modelling studies of HOMs and their
769 contributions to new particle formation and growth: comparison of boreal forest in Finland and a
770 polluted environment in China, *Atmos. Chem. Phys.*, 18, 11779-11791, 2018.

771

772 Quéléver, L. L. J., Kristensen, K., Normann Jensen, L., Rosati, B., Teiwes, R., Daellenbach, K. R.,
773 Peräkylä, O., Roldin, P., Bossi, R., Pedersen, H. B., Glasius, M., Bilde, M., and Ehn, M.: Effect of
774 temperature on the formation of highly oxygenated organic molecules (HOMs) from alpha-pinene
775 ozonolysis, *Atmos. Chem. Phys.*, 19, 7609-7625, <https://doi.org/10.5194/acp-19-7609-2019>, 2019.

776

777 Riccobono, F., Schobesberger, S., Scott, C., Dommen, J., Ortega, I., Rondo, L., Almeida, J.,
778 Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Downard, A., Dunne, E., Duplissy, J.,
779 Ehrhart, S., Flagan, R., Franchin, A., Hansel, A., Junninen, H., Kajos, M., Keskinen, H., Kupc, A.,
780 Kürten, A., Kvashin, A., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Nieminen, T.,

781 Onnela, A., Petäjä, T., Praplan, A., Santos, F., Schallhart, S., Seinfeld, J., Sipilä, M., Van Spracklen,
782 D., Stozhkov, Y., Stratmann, F., Tomé, A., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y.,
783 Vrtala, A., Wagner, P., Weingartner, E., Wex, H., Wimmer, D., Carslaw, K., Curtius, J., Donahue,
784 N., Kirkby, J., Kulmala, M., Worsnop, D., and Baltensperger, U.: Oxidation products of biogenic
785 emissions contribute to nucleation of atmospheric particles, *Science*, 344, 717-721, 2014.
786

787 Rissanen, M. P., Kurtén, T., Sipilä, M., Thornton, J. A., Kangasluoma, J., Sarnela, N., Junninen, H.,
788 Jørgensen, S., Schallhart, S., Kajos, M. K., Taipale, R., Springer, M., Mentel, T. F., Ruuskanen, T.,
789 Petäjä, T., Worsnop, D. R., Kjaergaard, H. G., and Ehn, M.: The formation of highly oxidized
790 multifunctional products in the ozonolysis of cyclohexene, *J. Am. Chem. Soc.*, 136, 15596-15606
791 2014.
792

793 Rissanen, M. P.: NO₂ Suppression of Autoxidation–Inhibition of Gas-Phase Highly Oxidized Dimer
794 Product Formation, *ACS Earth Space Chem.*, 2, 1211–1219,
795 <https://doi.org/10.1021/acsearthspacechem.8b00123>, 2018.
796

797 Rose, C., Zha, Q., Dada, L., Yan, C., Lehtipalo, K., Junninen, H., Mazon, S. B., Jokinen, T.,
798 Sarnela, N., Sipilä, M., Petäjä, T., Kerminen, V.-M., Bianchi, F., and Kulmala, M.: Observations of
799 biogenic ion-induced cluster formation in the atmosphere, *Sci. Adv.*, 4, eaar5218,
800 <https://doi.org/10.1126/sciadv.aar5218>, 2018.

801 Schobesberger, S., Junninen, H., Bianchi, F., Lönn, G., Ehn, M., Lehtipalo, K., Dommen, J.,
802 Ehrhart, S., Ortega, I. K., Franchin, A., Nieminen, T., Riccobono, F., Hutterli, M., Duplissy, J.,
803 Almeida, J., Amorim, A., Breitenlechner, M., Downard, A. J., Dunne, E. M., Flagan, R. C., Kajos,
804 M., Keskinen, H., Kirkby, J., Kupc, A., Kürten, A., Kurtén, T., Laaksonen, A., Mathot, S., Onnela,
805 A., Praplan, A. P., Rondo, L., Santos, F. D., Schallhart, S., Schnitzhofer, R., Sipilä, M., Tomé, A.,
806 Tsagkogeorgas, G., Vehkamäki, H., Wimmer, D., Baltensperger, U., Carslaw, K. S., Curtius, J.,
807 Hansel, A., Petäjä, T., Kulmala, M., Donahue, N. M., and Worsnop, D. R.: Molecular
808 understanding of atmospheric particle formation from sulfuric acid and large oxidized organic
809 molecules., *PNAS*, 110, 17223-17228, 2013.
810

811 Shi, J. P., Evans, D. E., Khan, A. A., and Harrison, R. M.: Sources and concentration of
812 nanoparticles (<10nm diameter) in the urban atmosphere, *Atmos. Environ.*, 35, 1193-1202, 2001.
813

814 Shi, Z., Vu, T., Kotthaus, S., Harrison, R.M., Grimmond, S., Yue, S., Zhu, T., Lee, J., Han, Y.,
815 Demuzere, M., Dunmore, R.E., Ren, L., Liu, D., Wang, Y., Wild, O., Allan, J., Acton, W.J.,
816 Barlow, J., Barratt, B., Beddows, D., Bloss, W.J., Calzolari, G., Carruthers, D., Carslaw, D.C., Chan,
817 Q., Chatzidiakou, L., Chen, Y., Crilley, L., Coe, H., Dai, T., Doherty, R., Duan, F., Fu, P., Ge, B.,
818 Ge, M., Guan, D., Hamilton, J.F., He, K., Heal, M., Heard, D., Hewitt, C.N., Hollaway, M., Hu, M.,
819 Ji, X. Jiang, R. Jones, M. Kalberer, F.J. Kelly, L. Kramer, B. Langford, C. Lin, A.C. Lewis, J. Li,
820 W. Li, D., Liu, H., Liu, J., Loh, M., Lu, K., Lucarelli, F., Mann, G., McFiggans, G., Miller, M.R.,
821 Mills, G., Monk, P., Nemitz, E., O'Connor, F., Ouyang, B., Palmer, P.I., Percival, C., Popoola, O.,
822 Reeves, C., Rickard, A.R., Shao, L., Shi, G., Spracklen, D., Stevenson, D., Sun, Y., Sun, Z., Tao, S.,
823 Tong, S., Wang, Q., Wang, W., Wang, X., Wang, X., Wang, Z., Wei, L., Whalley, L., Wu, X., Wu,
824 Z., Xie, P., Yang, F., Zhang, Q., Zhang, Y., Zhang, Y. and Zheng, M.: In-depth study of air
825 pollution sources and processes within Beijing and its surrounding region (APHH-Beijing), *Atmos.*
826 *Chem. Phys.*, 19, 7519-7546, 2019.
827

828 Stolzenburg, D., Fischer, L., Vogel, A., Heinritzi, M., Schervish, M., Simon, M., Wagner, A.,
829 Dada, L., Ahonen, L., Amorim, A., Baccharini, A., Bauer, P., Baumgartner, B., Bergen, A.,
830 Bianchi, F., Breitenlechner, M., Brilke, S., Buenrostro Mazon, S., Chen, D., Dias, A., Draper, D.,
831 Duplissy, J., El Haddad, I., Finkenzeller, H., Frege, C., Fuchs, C., Garmash, O., Gordon, H., He, X.,
832 Helm, J., Hofbauer, V., Hoyle, C., Kim, C., Kirkby, J., Kontkanen, J., Kürten, A., Lampilahti, J.,

833 Lawler, M., Lehtipalo, K., Leiminger, M., Mai, H., Mathot, S., Mentler, B., Molteni, U., Nie, W.,
834 Nieminen, T., Nowak, J., Ojdanic, A., Onnela, A., Passananti, M., Petäjä, T., Quéléver, L.,
835 Rissanen, M., Sarnela, N., Schallhart, S., Tauber, C., Tomé, A., Wagner, R., Wang, M., Weitz, L.,
836 Wimmer, D., Xiao, M., Yan, C., Ye, P., Zha, Q., Baltensperger, U., Curtius, J., Dommen, J., Flagan,
837 R., Kulmala, M., Smith, J., Worsnop, D., Hansel, A., Donahue, N., Winkler, P., Rapid growth of
838 organic aerosol nanoparticles over a wide tropospheric temperature range, *PNAS*, 115, 9122-9127,
839 2018.

840

841 Tammet, H.: Size and mobility of nanometer particles, clusters and ions, *J. Aerosol Sci.*, 26, 459-
842 475, 1995.

843

844 Tomasi, C., Fuzzi, S., Kokhanovsky, A.: *Atmospheric Aerosols: Life Cycles and Effects on Air*
845 *Quality and Climate*, Wiley, 2016.

846

847 Tröstl, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Ahlm, L., Frege, C.,
848 Bianchi, F., Wagner, R., Simon, M., Lehtipalo, K., Williamson, C., Craven, J. S., Duplissy, J.,
849 Adamov, A., Almeida, J., Bernhammer, A. K., Breitenlechner, M., Brilke, S., Dias, A., Ehrhart, S.,
850 Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Gysel, M., Hansel, A., Hoyle, C. R., Jokinen, T.,
851 Junninen, H., Kangasluoma, J., Keskinen, H., Kim, J., Krapf, M., Kürten, A., Laaksonen, A.,
852 Lawler, M., Leiminger, M., Mathot, S., Möhler, O., Nieminen, T., Onnela, A., Petäjä, T., Piel, F.
853 M., Miettinen, P., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Sengupta, K., Sipilä,
854 M., Smith, J. N., Steiner, G., Tomé, A., Virtanen, A., Wagner, A. C., Weingartner, E., Wimmer, D.,
855 Winkler, P. M., Ye, P., Carslaw, K. S., Curtius, J., Dommen, J., Kirkby, J., Kulmala, M., Riipinen,
856 I., Worsnop, D. R., Donahue, N. M., and Baltensperger, U.: The role of low-volatility organic
857 compounds in initial particle growth in the atmosphere, *Nature*, 533, 527-531,
858 <https://doi.org/10.1038/nature18271>, 2016.

859

860 Wang, S., Wu, R., Berndt, T., Ehn, M., and Wang, L.: Formation of Highly Oxidized Radicals and
861 Multifunctional Products from the Atmospheric Oxidation of Alkylbenzenes, *Environ. Sci. Techn.*,
862 51, 8442-8449, 2017.

863

864 Wang, Z., Wu, Z., Yue, D., Shang, D., Guo, S., Sun, J., Ding, A., Wang, L., Jiang, J., Guo, H., Gao,
865 J., Cheung, H. C., Morawska, L., Keywood, M., and Hu, M.: New particle formation in China:
866 Current knowledge and further directions, *Sci. Tot. Environ.*, 577, 258-266, 2016.

867

868 Wiedensohler, A., Cheng, Y. F., Nowak, A., Wehner, B., Achtelt, P., Berghof, M., Birmili, W., Wu,
869 Z. J., Hu, M., Zhu, T., Takegawa, N., Kita, K., Kondo, Y., Lou, S. R., Hofeumahaas, A., Holland,
870 F., Wahner, A., Gunthe, S. S., Rose, D., Su, H., and Pöschl, U.: Rapid aerosol particle growth and
871 increase of cloud condensation nucleus activity by secondary aerosol formation and condensation:
872 A case study for regional air pollution in northeastern China, *Journal of Geophysical Research*
873 *Atmospheres*, 114, 1–13, <https://doi.org/10.1029/2008JD010884>, 2009.

874

875 Wu, Z., Hu, M., Lin, P., Liu, S., Wehner, B., and Wiedensohler, A.: Particle number size
876 distribution in the urban atmosphere of Beijing, China, *Atmos. Environ.*, 42, 7967-7980, 2008.

877

878 Wu, Z., Ma, N., Größ, J., Kecorius, S., Lu, K., Shang, D., Wang, Y., Wu, Y., Zeng, L., Hu, M.,
879 Wiedensohler, A., and Zhang, Y.: Thermodynamic properties of nanoparticles during new particle
880 formation events in the atmosphere of North China Plain, *Atmos. Res.*, 188, 55-63, 2017.

881

882

883

884 Wu, Z. J., Hu, M., Liu, S., Wehner, B., Bauer, S., Ssling, a. M., Wiedensohler, a., Petäjä, T., Dal
885 Maso, M., and Kulmala, M.: New particle formation in Beijing, China: Statistical analysis of a 1-
886 year data set, *J. Geophys. Res. Atmospheres*, 112, D09209, <https://doi.org/10.1029/2006JD007406>,
887 2007.
888

889 Xiong, F., McAvey, K. M., Pratt, K. A., Groff, C. J., Hostetler, M. A., Lipton, M. A., Starn, T. K.,
890 Seeley, J. V., Bertman, S. B., Teng, A. P., Crouse, J. D., Nguyen, T. B., Wennberg, P. O., Misztal,
891 P. K., Goldstein, A. H., Guenther, A. B., Koss, A. R., Olson, K. F., De Gouw, J. A., Baumann, K.,
892 Edgerton, E. S., Feiner, P. A., Zhang, L., Miller, D. O., Brune, W. H., and Shepson, P. B.:
893 Observation of isoprene hydroxynitrates in the Southeastern United States and implications for the
894 fate of NO_x, *Atmos. Chem. Phys.*, 15, 11257-11272, 2015 .
895

896 Yan, C., Nie, W., Äijälä, M., Rissanen, M. P., Canagaratna, M. R., Massoli, P., Junninen, H.,
897 Jokinen, T., Sarnela, N., Häme, S. A. K., Schobesberger, S., Canonaco, F., Yao, L., Prévôt, A. S.
898 H., Petäjä, T., Kulmala, M., Sipilä, M., Worsnop, D. R., and Ehn, M.: Source characterization of
899 highly oxidized multifunctional compounds in a boreal forest environment using positive matrix
900 factorization, *Atmos. Chem. Phys.*, 16, 12715-12731, 2016.
901

902 Yao, L., Garmash, O., Bianchi, F., Zheng, J., Yan, C., Kontkanen, J., Junninen, H., Mazon, S. B.,
903 Ehn, M., Paasonen, P., Sipilä, M., Wang, M., Wang, X., Xiao, S., Chen, H., Lu, Y., Zhang, B.,
904 Wang, D., Fu, Q., Geng, F., Li, L., Wang, H., Qiao, L., Yang, X., Chen, J., Kerminen,
905 V. M., Petäjä, T., Worsnop, D. R., Kulmala, M., and Wang, L.: Atmospheric new particle formation
906 from sulfuric acid and amines in a Chinese megacity, *Science*, 361, 278-281, 2018.
907

908 Yu, F. and Luo, G.: Simulation of particle size distribution with a global aerosol model:
909 Contribution of nucleation to aerosol and CCN number concentrations, *Atmospheric Chemistry and
910 Physics*, 9, 7691-7710, 2009.
911

912 Yue, D. L., Hu, M., Zhang, R. Y., Wu, Z. J., Su, H., Wang, Z. B., Peng, J. F., He, L. Y., Huang, X.
913 F., Gong, Y. G., and Wiedensohler, A.: Potential contribution of new particle formation to cloud
914 condensation nuclei in Beijing, *Atmos. Environ.*, 45, 6070-6077, 2011.
915

916 Zhao, Y., Wingen, L. M., Perraud, V., Greaves, J., and Finlayson-Pitts, B. J.: Role of the reaction of
917 stabilized Criegee intermediates with peroxy radicals in particle formation and growth in air, *Phys.
918 Chem. Chem. Phys.*, 17, 12500-12514, 2015.
919

920 **FIGURE LEGENDS:**

921

922 **Figure 1** Oxidation state of carbon calculated as two times the oxygen to carbon ratio minus the
923 hydrogen to carbon ratio against carbon number for (colored) individual ions and (blue
924 circles) signal weighted average for each carbon number. Area and colour are both
925 proportional to the peak area for each ion
926

927 **Figure 2** Mass defect plot of fitted mass spectral peaks between 100-600mass units on (a) 10:30
928 – 12:00 23/06/2017, a non nucleation day, and (b) 10:30 -12:00 25/06/2017, a
929 nucleation day. Mass defect can be defined as the mass - integer mass. The size of point
930 is proportional to the signal intensity. As ^1H has a positive mass defect (1.007276 Da),
931 the upward trend along the horizontal indicates increasing carbon chain length, and
932 differences at similar masses are due to increasing oxygen functionality, clustering with
933 species such as sulfuric acid (negative mass defect) and ammonia (positive mass
934 defect), as ^{16}O and ^{32}S have negative mass defects (15.9949 and 31.9721 Da
935 respectively), while ^{14}N has a positive mass defect at 14.0031 Da. The two large peaks
936 seen at 201 and 288 m/Q are the nitrophenol-nitrate cluster and a $\text{C}_5\text{H}_{10}\text{N}_2\text{O}_8$ -nitrate
937 cluster respectively.
938

939 **Figure 3** Summed time series of the normalised signals of (A) all non-nitrogen containing HOMs
940 and all organonitrates identified, (B) C5, C10 and C20 components, assumed to be
941 dominated by isoprene, monoterpene monomer and monoterpene dimers, signal for C20
942 multiplied 50 times to fit scale, and (C) summed C6 - C9 components, and summed C11
943 - C18 components, assumed to be dominated by alkylbenzenes and other larger
944 components respectively.
945

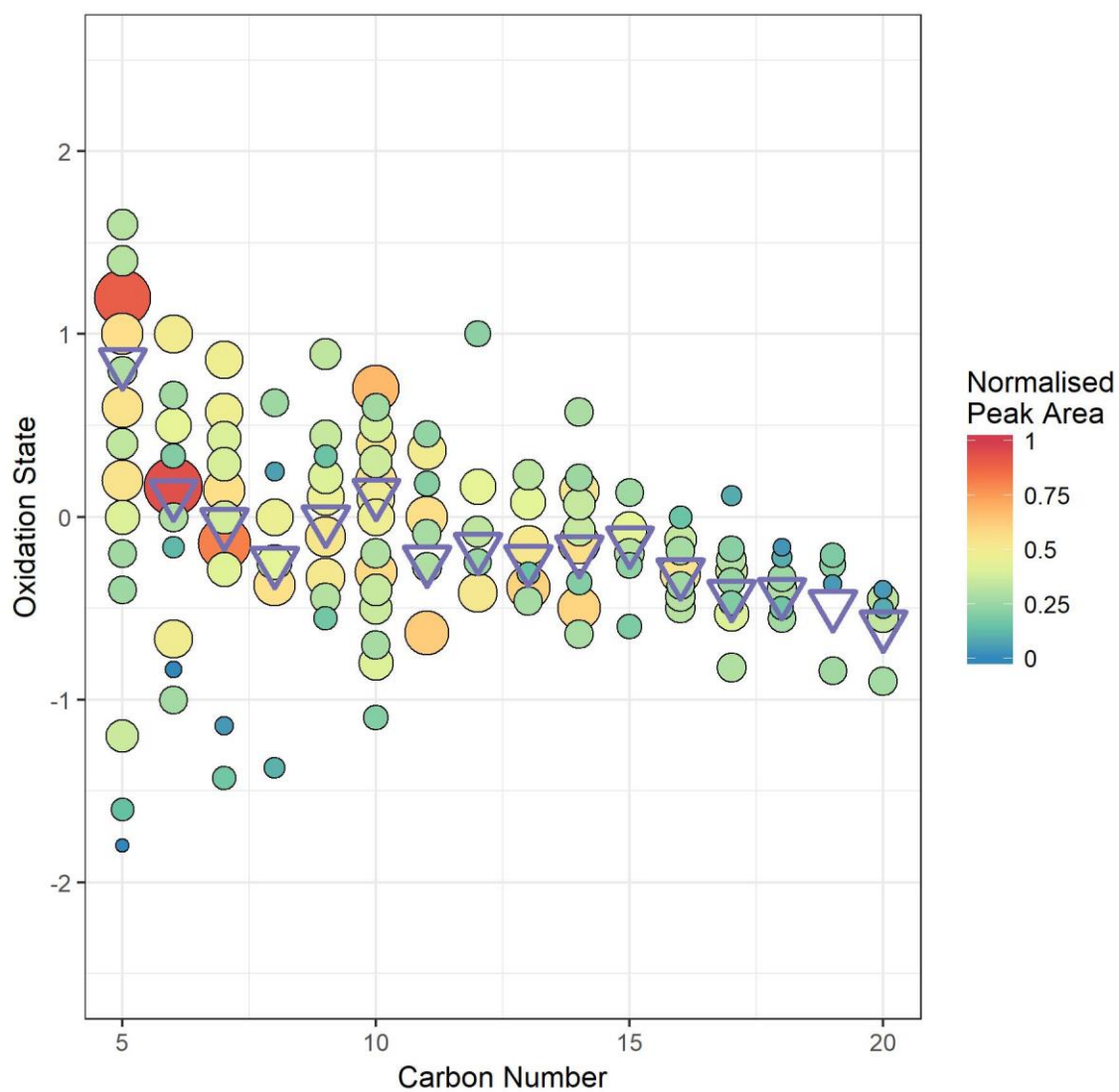
946 **Figure 4** Time series for the whole sampling campaign for the concentrations of (left axis) VOCs
947 as measured by PTR-ToF and (right axis) a selected HOM product associated with that
948 precursor.
949

950 **Figure 5** Normalised unit mass NO_3 - CI-APi-ToF signal intensity on 24/06/2017 (A) and
951 25/06/2017 (B). Each individual unit mass was normalised to a maximum of 1. Each
952 period is normalised separately so the individual signal maxima on each day are visible.
953 The graph is plotted between 200-600 mass units, with every 10 mass units averaged for
954 simplicity. On the secondary axis is plotted PSM data, both total particle count >1.30
955 nm (black trace) and total clusters between 1.30 and 1.84 nm (blue trace). Data is
956 plotted at 1 hour time resolution.
957

958 **Figure 6** SMPS + PSM contour plot for two nucleation days on 24/06/2017 and 25/06/2017. Data
959 in bottom panel is from the PSM instrument, top panel from NanoSMPS, units in colour
960 bar are $\log_{10}(\text{dN}/\log D_p)$ for N in cm^{-3} . Points signify normalised sulfuric acid
961 concentration (right axis) as measured by CI-APi-ToF.
962

963

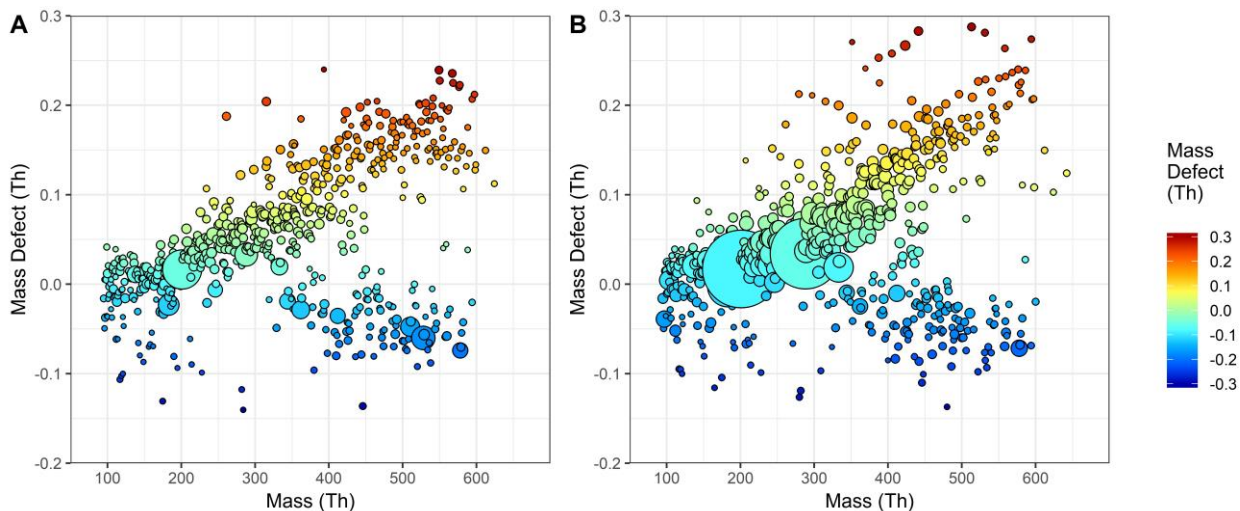
964



965

966 **Figure 1.** Oxidation state of carbon plotted against carbon number for (colored) individual ions and
967 (purple triangles) signal weighted average for each carbon number. Area and colour are both
968 proportional to the peak area for each ion.

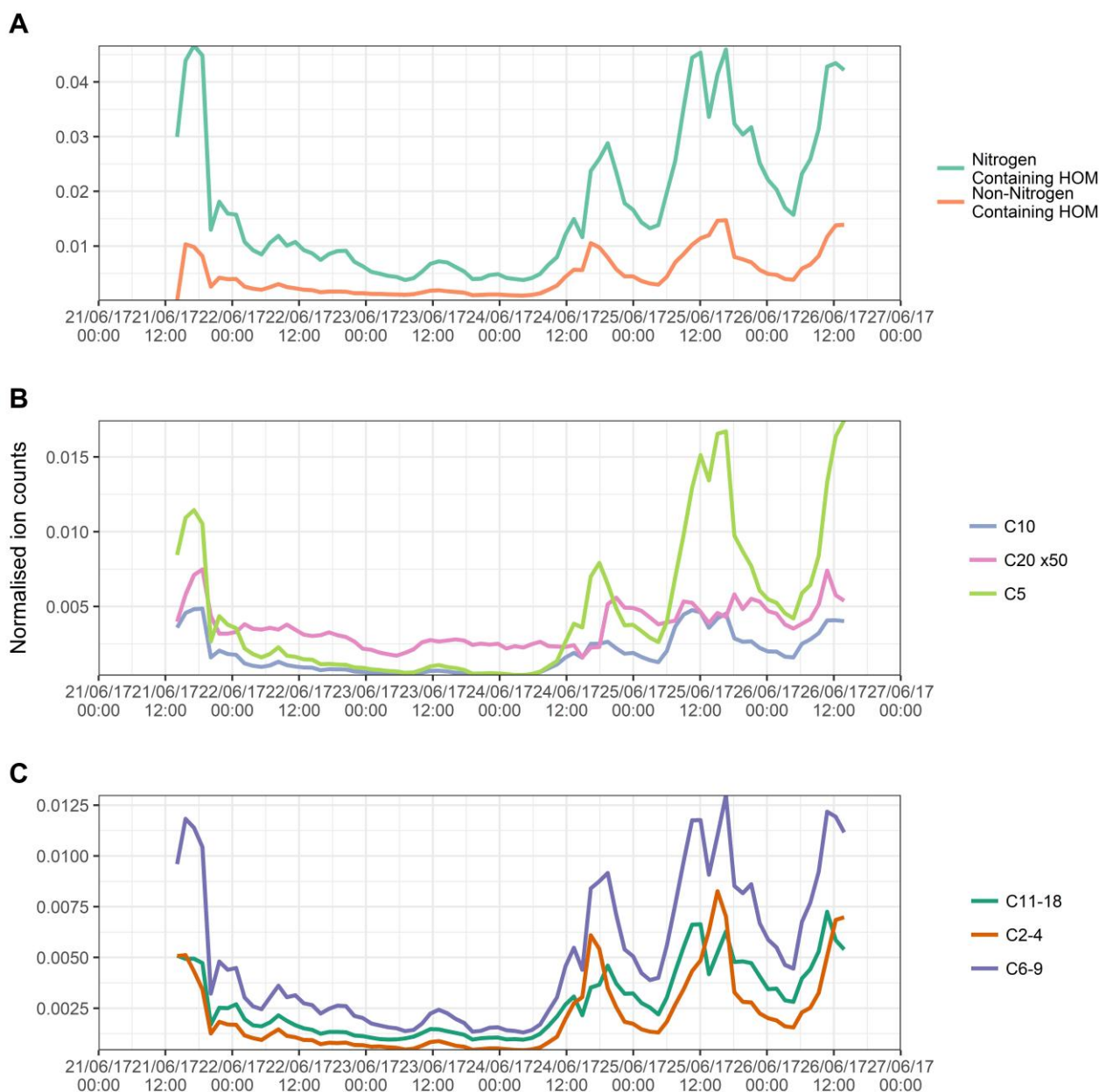
969



971

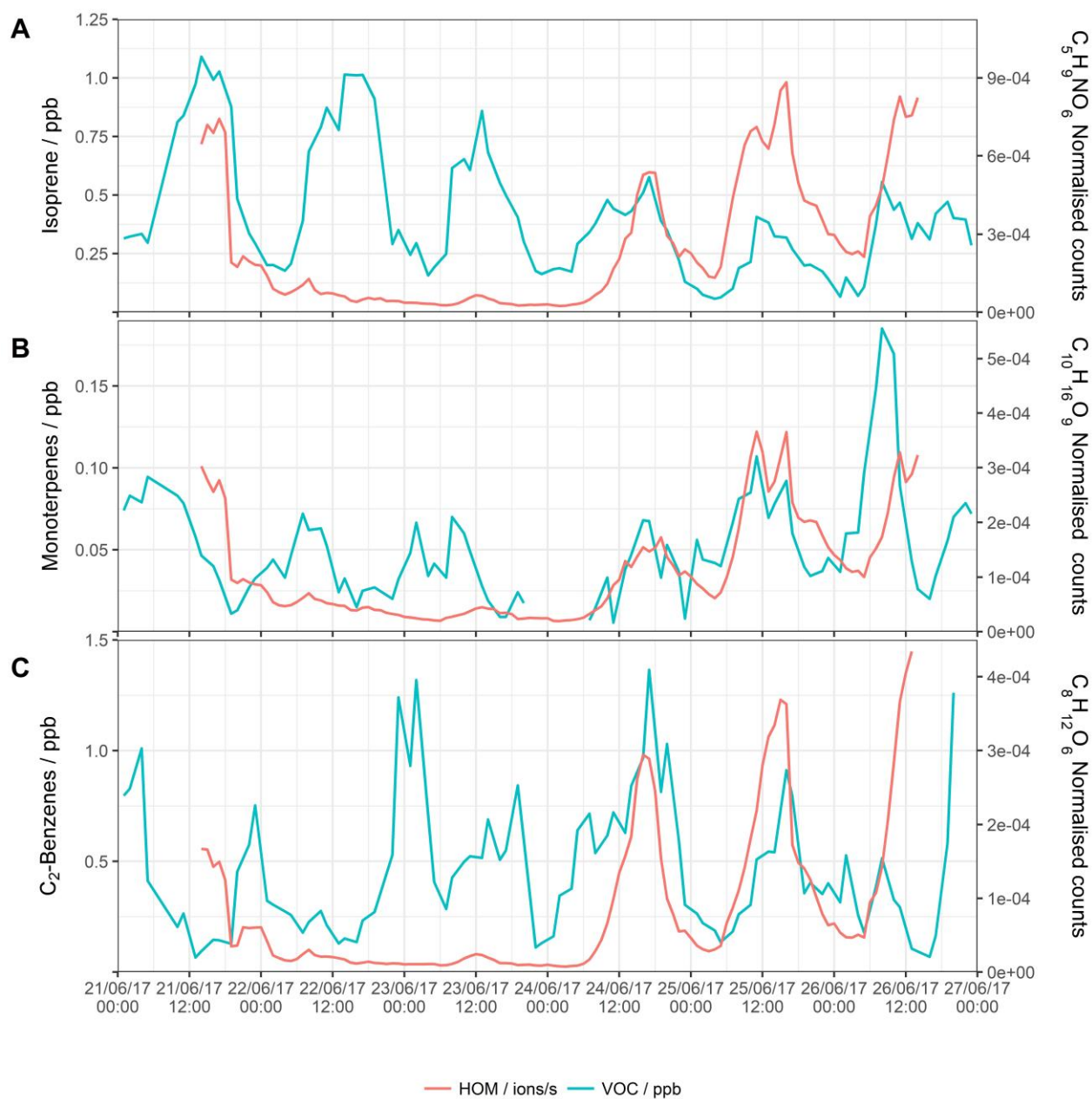
972 **Figure 2.** Mass defect plot of fitted mass spectral peaks between 100-600 mass units on (a) 10:30 –
 973 12:00 23/06/2017, a non nucleation day, and (b) 10:30 -12:00 25/06/2017, a nucleation day. Mass
 974 defect can be defined as the mass - integer mass. The size of point is proportional to the signal
 975 intensity. As ^1H has a positive mass defect (1.007276 Da), the upward trend along the horizontal
 976 indicates increasing carbon chain length, and differences at similar masses are due to increasing
 977 oxygen functionality, clustering with species such as sulfuric acid (negative mass defect) and
 978 ammonia (positive mass defect), as ^{16}O and ^{32}S have negative mass defects (15.9949 and 31.9721 Da
 979 respectively), while ^{14}N has a positive mass defect at 14.0031 Da. The two large peaks seen at 201
 980 and 288 m/Q are the nitrophenol-nitrate cluster and a $\text{C}_5\text{H}_{10}\text{N}_2\text{O}_8$ -nitrate cluster respectively.

981



982

983 **Figure 3.** Summed time series of the normalised signals of (A) all non-nitrogen containing HOMs
 984 and all organonitrates identified, (B) C5, C10 and C20 components, assumed to be dominated by
 985 isoprene, monoterpene monomer and monoterpene dimers, signal for C20 multiplied 50 times to fit
 986 scale, and (C) summed C6 - C9 components, and summed C11 - C18 components, assumed to be
 987 dominated by alkylbenzenes and other larger components respectively.



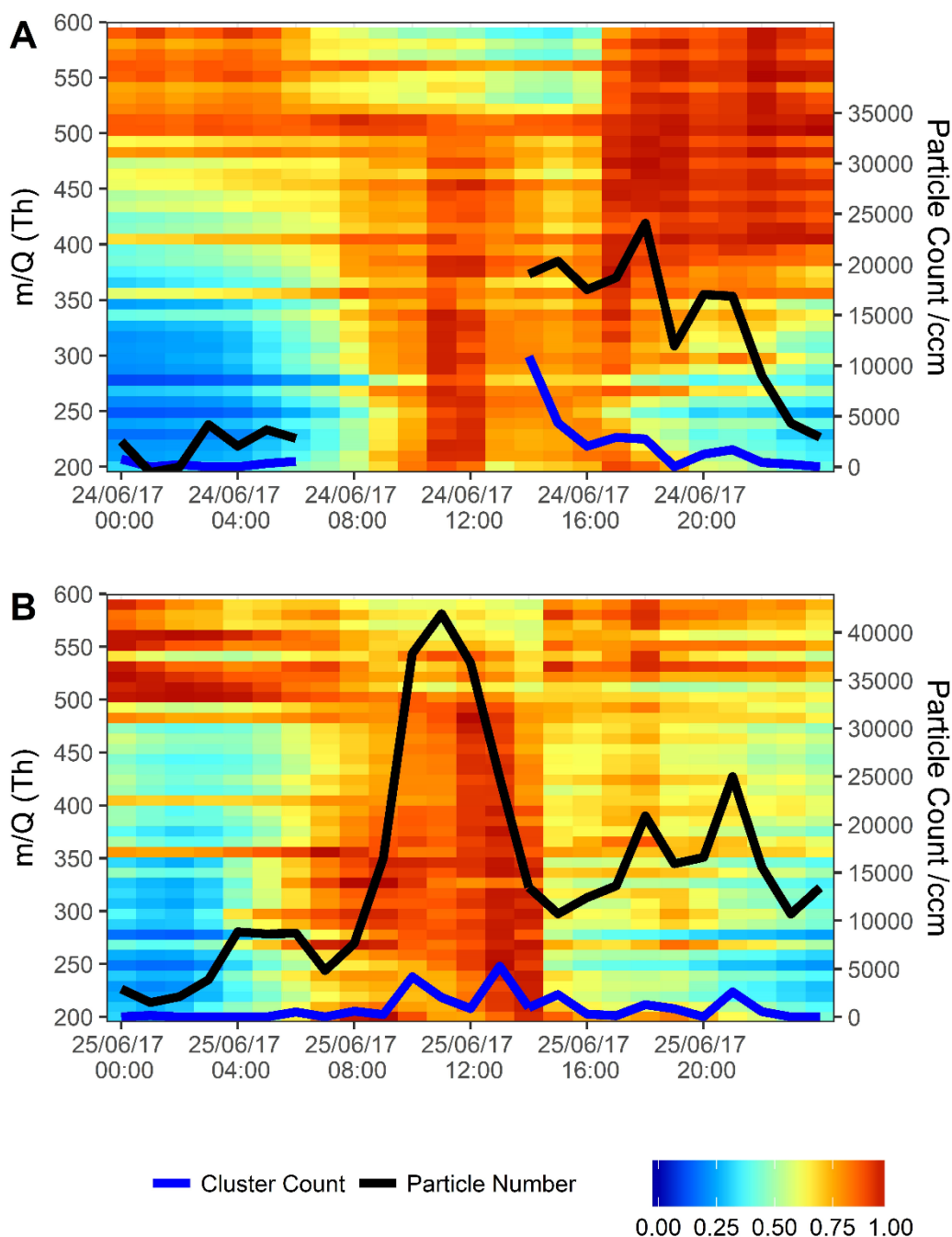
988

989 **Figure 4.** Time series for the whole sampling campaign for the concentrations of (left axis) VOCs as
 990 measured by PTR-ToF and (right axis) a selected HOM product associated with that precursor.

991

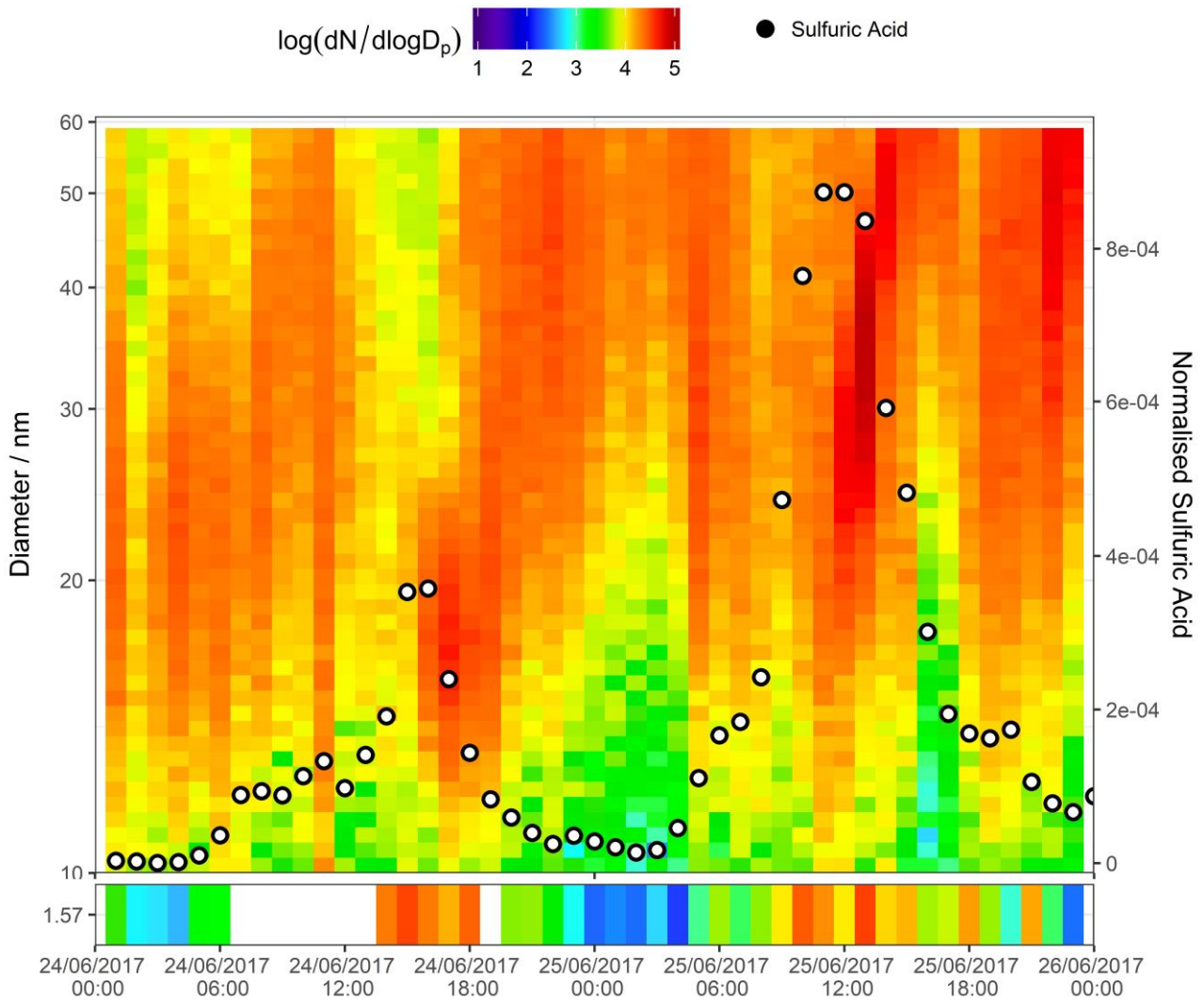
992

993



994

995 **Figure 5.** Normalised unit mass NO_3^- CI-APi-ToF signal intensity on 24/06/2017 (A) and 25/06/2017
 996 (B). Each individual unit mass was normalised to a maximum of 1. Each period is normalised
 997 separately so the individual signal maxima on each day are visible. The graph is plotted between 200-
 998 600 mass units, with every 10 mass units averaged for simplicity. On the secondary axis is plotted
 999 PSM data, both total particle count >1.30 nm (black trace) and total clusters between 1.30 and 1.84
 1000 nm (blue trace). Data is plotted at 1 hour time resolution.



1001

1002 **Figure 6.** SMPS + PSM contour plot for two nucleation days on 24/06/2017 and 25/06/2017. Data in
 1003 bottom panel is from the PSM instrument, top panel from NanoSMPS, units in colour bar are \log_{10}
 1004 $(dN/d\log D_p)$ for N in cm^{-3} . Points signify normalised sulfuric acid concentration (right axis) as
 1005 measured by CI-APi-ToF.

1006
 1007

1008
 1009

1010
 1011

1012