



1 **Exploiting multi-wavelength aerosol absorption coefficients in a**
2 **multi-time source apportionment study to retrieve source-**
3 **dependent absorption parameters**

4 Alice C. Forello¹, Vera Bernardoni¹, Giulia Calzolai², Franco Lucarelli², Dario Massabò³, Silvia
5 Nava², Rosaria E. Pileci^{1,a}, Paolo Prati³, Sara Valentini¹, Gianluigi Valli¹, Roberta Vecchi¹

6 ¹Department of Physics, Università degli Studi di Milano and National Institute of Nuclear Physics INFN-Milan, via
7 Celoria 16, Milan, 20133, Italy

8 ²Department of Physics and Astronomy, Università di Firenze and National Institute of Nuclear Physics INFN-Florence,
9 via G. Sansone 1, Sesto Fiorentino, 50019, Italy

10 ³Department of Physics, Università degli Studi di Genova and National Institute of Nuclear Physics INFN- Genoa, via
11 Dodecaneso 33, Genoa, 16146, Italy

12 ^anow at: Laboratory of Atmospheric Chemistry (LAC), Paul Scherrer Institut (PSI), Forschungsstrasse 111, Villigen,
13 5232, Switzerland

14
15 *Correspondence to:* Roberta Vecchi (roberta.vecchi@unimi.it)

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18 **Abstract.** In this paper, a new methodology coupling aerosol optical and chemical parameters in the same source
19 apportionment study is reported. This approach gives additional relevant information such as estimates for the
20 atmospheric Ångström Absorption Exponent (α) of the sources and Mass Absorption Coefficient (MAC) for fossil fuel
21 emissions at different wavelengths.

22 A multi-time source apportionment study using Multilinear Engine ME-2 was performed on a PM10 dataset with different
23 time resolution (24 hours, 12 hours, and 1 hour) collected during two different seasons in Milan (Italy) in 2016. Samples
24 were optically analysed to retrieve the aerosol absorption coefficient b_{ap} (in Mm^{-1}) at four wavelengths ($\lambda=405$ nm, 532
25 nm, 635 nm and 780 nm) and chemically characterised for elements, ions, levoglucosan, and carbonaceous components.
26 Time-resolved chemically speciated data were coupled with b_{ap} multi-wavelength measurements and introduced as input
27 data in the multi-time receptor model; this approach was proven to strengthen the identification of sources being
28 particularly useful when important chemical markers (e.g. levoglucosan, elemental carbon, ...) are not available. The



29 final solution consisted in 8 factors (nitrate, sulphate, resuspended dust, biomass burning, construction works, traffic,
30 industry, aged sea salt); the implemented constraints led to a better physical description of factors and the bootstrap
31 analysis supported the goodness of the solution. As for b_{ap} apportionment, consistently to what expected, the two factors
32 assigned to biomass burning and traffic were the main contributors to aerosol absorption in atmosphere. A relevant feature
33 of the approach proposed in this work is the possibility of retrieving many other information about optical parameters;
34 for example, opposite to the more traditional approach used by optical source apportionment models, here we obtained
35 the atmospheric Ångström Absorption Exponent (α) of the sources (α biomass burning = 1.83 and α fossil fuels = 0.80),
36 without any a priori assumption. In addition, an estimate for the Mass Absorption Cross section (MAC) for fossil fuel
37 emissions at four wavelengths was obtained and found to be consistent with literature ranges.

38

39 1. Introduction

40 Atmospheric aerosol impacts both on local and global scale causing adverse health effects (Pope and Dockery, 2006),
41 decreasing visibility (Watson, 2002) and influencing the climate (IPCC, 2013). To face these issues an accurate
42 knowledge of the aerosol emission sources is mandatory.

43 At the state of the art, multivariate receptor models are considered a robust approach (Belis et al., 2015) to carry out
44 source apportionment studies of atmospheric aerosol and the Positive Matrix Factorization (PMF) (Paatero and Tapper,
45 1994) has become one of the most widely used receptor models (Hopke, 2016). In the late 1990s the Multilinear Engine
46 (ME2) was implemented, a very flexible algorithm to solve multilinear and quasi-multilinear problems (Paatero, 1999).
47 This algorithm introduced the possibility to write scripts and implement advanced receptor modelling approaches; one
48 example is the multi-time model, developed for the first time by Zhou et al. (2004), which uses each experimental data in
49 its original time schedule as model input. This approach is very useful in measurement campaigns when instrumentation
50 with different time resolution (minutes, hours or days) is available; indeed, with the multi-time approach, high time
51 resolution data can be exploited without averaging them over the longest sampling interval. Nevertheless, source
52 apportionment studies carried out by multi-time model are still scarce in the literature (Zhou et al., 2004; Ogulei et al.,
53 2005; Kuo et al., 2014; Liao et al., 2015; Crespi et al., 2016; Sofowote et al., 2018).

54 The combination of time-resolved chemically speciated data with the information obtained from instrumentation
55 measuring aerosol optical properties at different wavelengths (e.g. the absorption coefficient b_{ap}) is suggested as one of
56 the future investigations of receptor modelling (Hopke, 2016); however, to the best of our knowledge, very few attempts
57 in this direction have been done (recently e.g. Xie et al., 2019). Wang et al. (2011, 2012) introduced in a source
58 apportionment study the Delta-C (Delta-C = $BC@370\text{ nm} - BC@880\text{ nm}$ from aethalometer measurements) as an input



59 variable in a source apportionment study and found that Delta-C was very useful in separating traffic from wood burning
60 source contributions.

61 The wavelength dependence of b_{ap} can be empirically considered proportional to $\lambda^{-\alpha}$, where α is the Ångström
62 Absorption Exponent; α depends on particles composition and size, and it is a useful parameter to gain information about
63 particles type in atmosphere (see e.g. Yang et al., 2009). Among PM components, black carbon (BC) is the main
64 responsible for light absorption in atmosphere; in fact, it is considered the main PM contributor to global warming and
65 the second most important anthropogenic contributor after CO₂ (Bond et al., 2013). Black carbon refers to a fraction of
66 the carbonaceous aerosol that shares peculiar features about microstructure, morphology, thermal stability, solubility, and
67 light absorption (Petzold et al., 2013); in particular, it is characterised by a wavelength-independent imaginary part of the
68 refractive index over visible and near-visible regions. In the last decade, experimental studies evidenced also the role of
69 another absorbing component i.e. brown carbon (BrC), referred to as light-absorbing organic matter of various origins
70 with increasing absorption towards lower wavelengths, especially in the UV region (Andreae and Gelencsér, 2006). BrC
71 is an aerosol component that also affects the elemental vs. organic carbon correct separation when using thermal-optical
72 methods as recently outlined by Massabò et al. (2016).

73 Source apportionment optical models based only on multi-wavelength measurements of b_{ap} are available in the literature,
74 i.e. the widespread Aethalometer model (Sandradewi et al., 2008) and the more recent Multi-Wavelength Absorption
75 Analyzer (MWAA) model (Massabò et al., 2015; Bernardoni et al., 2017b). Briefly, these models allow to estimate the
76 contribution of sources to aerosol absorption in atmosphere exploiting their different dependence on λ (different α). As a
77 step forward, MWAA provides the b_{ap} apportionment in relation to both the sources and the components (i.e. BC and
78 BrC) and gives also an estimate for α of BrC. Source apportionment optical models usually assume two contributors to
79 b_{ap} , namely fossil fuels combustion and wood burning (only few exceptions are present in the literature, e.g. Fialho et al.,
80 2005). In most cases this assumption is well founded, except in presence of episodic events that give a not negligible
81 contribution to aerosol absorption in atmosphere, such as the transport of mineral dust from the Saharan desert (Fuzzi et
82 al., 2015). Moreover, the above-mentioned models need a priori assumption about α of the sources; this is the most critical
83 step, since α depends on the kind of fuel, burning conditions and aging processes in the atmosphere and wide ranges for
84 α are reported in literature (Sandradewi et al., 2008). Without accurate determination of source-specific atmospheric α
85 (for example exploiting the information derived from source apportionment using ¹⁴C measurements), the applicability
86 of models based on optical determination is questionable (Bernardoni et al., 2017b; Massabò et al., 2015; Zotter et al.,
87 2017). Moreover, the generally accepted assumption of $\alpha=1$ for fossil fuels and BC, that is derived from the theory of
88 absorption of spherical particles in the Rayleigh regime (Seinfeld and Pandis, 2006), might not always be valid in
89 atmosphere due to aerosol aging processes (Liu et al., 2018).



90 In this work, in the frame of a source apportionment study based on multi-time receptor modelling, optical and chemical
91 data were coupled to explore the possibility of retrieving a multi- λ apportionment of b_{ap} with no need of a-priori
92 assumptions on the contributing sources. Opposite, with this approach source-dependent α values can be provided as
93 output. Moreover, the multi- λ apportionment of b_{ap} in each source allowed to estimate MAC values at different
94 wavelengths, exploiting the well-known relation $EBC=b_{ap}(\lambda)/MAC(\lambda)$ (Bond and Bergstrom, 2006) and considering the
95 apportioned concentrations of elemental carbon (EC) as a proxy for BC. It is noteworthy that the evaluation of
96 atmospheric MAC values is not trivial due to the possible presence of absorbing components different from BC (e.g.
97 contribution from BrC, especially at lower wavelengths).

98 The original approach proposed in this work shows that coupling the chemical and optical information in a receptor
99 modelling process is particularly advantageous because: (1) strengthens the source identification, that is particularly
100 useful when relevant chemical tracers (e.g. levoglucosan, EC, ...) are not available; (2) gives estimates for source-specific
101 atmospheric Ångström Absorption Exponent (α) which are typically assumed a-priori in optical apportionment models;
102 (3) assesses MAC values at different wavelengths for specific sources.

103

104 2. Material and methods

105 2.1 Site description and aerosol sampling

106 Two measurement campaigns were performed during summertime (June-July) and wintertime (November-December)
107 2016 in Milan (Italy). Milan is the largest city (more than 1 million inhabitants, doubled by commuters everyday) of the
108 Po Valley, a very well-known hot-spot pollution area in Europe due to both large emissions from a variety of sources (i.e.
109 traffic, industry, domestic heating, energy production plants, and agriculture) and low atmospheric dispersion conditions
110 (e.g. Vecchi et al., 2007 and 2019; Perrone et al., 2012; Bigi and Ghermandi, 2014; Perrino et al., 2014).

111 The sampling site is representative of the urban background and it is situated at about 10 meters above the ground, on the
112 roof of the Physics Department of the University of Milan, less than 4 km far from the city centre (Vecchi et al., 2009).

113 It is important to note that during the sampling campaigns, a large building site was in activity next to the monitoring
114 station.

115 Aerosol sampling was carried out using instrumentation with different time-resolution. Low time resolution PM10 data,
116 with a sampling duration of 24 and 12 hours during summertime (20 June-22 July 2016) and wintertime (21 November-
117 22 December 2016), respectively, were collected in parallel on PTFE (Whatman, 47 mm diameter) and pre-fired (700 °C,
118 1 hour) quartz-fibre (Pall, 2500QAO-UP, 47 mm diameter) filters. Low volume samplers with EPA PM10 inlet operating
119 at 1 m³ h⁻¹ were used. High time resolution data were collected during shorter periods (11 July-18 July and 21 November-
120 28 November 2016) by a streaker sampler (D'Alessandro et al., 2003). Shortly, the streaker sampler collects the fine and



121 coarse PM fractions (particles with aerodynamic diameter $d_{ae} < 2.5 \mu\text{m}$, and $2.5 < d_{ae} < 10 \mu\text{m}$, respectively) with hourly
122 resolution. Particles with $d_{ae} > 10 \mu\text{m}$ impact on the first stage and are discarded; the coarse fraction deposits on the second
123 stage, consisting of a Kapton foil; finally, the fine fraction is collected on a polycarbonate filter. The two collecting
124 supports are kept in rotation with an angular speed of about 1.8°h^{-1} to produce a circular continuous deposit on both
125 stages.

126 Meteorological data were available at a monitoring station belonging to the regional environmental agency (ARPA
127 Lombardia) which is less than 1 km far away.

128

129 *2.2 PM mass concentration and chemical characterisation*

130 PM₁₀ mass concentration was determined on PTFE filters by gravimetric technique. Weighing was performed by an
131 analytical balance (Mettler, model UMT5, 1 μg sensitivity) after a 24 hours conditioning period in an air-controlled room
132 as for temperature ($20 \pm 1^\circ\text{C}$) and relative humidity ($50 \pm 3\%$) (Vecchi et al., 2004).

133 These filters were then analysed by Energy Dispersive X-Ray Fluorescence (ED-XRF) analysis to obtain the elemental
134 composition (details on the procedure can be found in Vecchi et al., 2004). For most elements and samples, concentrations
135 were characterised by relative uncertainties in the range 7-20 % (higher uncertainties for elements with concentrations
136 next to MDL) and Minimum Detection Limits (MDL) of 0.9-30 ng m^{-3} with the above mentioned sampling conditions.

137 Quartz-fibre filters were punched and analysed to detect levoglucosan, sulphate, nitrate, and carbonaceous components.
138 For each filter, one punch (1.5cm^2) was extracted by sonication (1 h) using 5 ml ultrapure Milli-Q water; this extract was

139 analysed to measure both levoglucosan and inorganic anions concentrations. Levoglucosan concentration was determined
140 by High-Performance Anion Exchange Chromatography coupled with Pulsed Amperometric Detection (HPAEC-PAD)

141 (Piazzalunga et al., 2010) only in winter samples. Indeed, as already pointed out by other studies at the same sampling
142 site (Bernardoni et al., 2011) and as routinely measured at monitoring stations in Milan by the Regional Environmental

143 Agency (private communication), levoglucosan concentrations during summertime are lower than the MDL of the
144 technique (about 6ng m^{-3}), due to both lower emissions (no influence of residential heating and negligible impact from

145 other sources) and higher OH levels in the atmosphere depleting molecular markers concentrations (Robinson et al., 2006;
146 Hennigan et al., 2010). Uncertainties on levoglucosan concentration were about 11 %. The measurement of main water-

147 soluble inorganic anions (SO_4^{2-} and NO_3^-) was performed by Ion Chromatography (IC) with MDL of 25 and 50 ng m^{-3}
148 with summertime and wintertime sampling conditions, respectively, and uncertainties of about 10 %. Unfortunately, due

149 to technical problems no data on ammonium were available. Details on the analytical procedure for IC analysis are
150 reported in Piazzalunga et al. (2013). Another punch (1.0cm^2) of each filter was analysed by Thermal Optical

151 Transmittance analysis (TOT, Sunset Inc., NIOSH-870 protocol) (Piazzalunga et al., 2011) in order to assess organic and



152 elemental carbon (OC and EC) concentrations. MDL was 75 and 150 ng m⁻³ with summertime and wintertime sampling
153 conditions, respectively, and uncertainties were about 10-15 %.

154 High time resolution samples were analysed to obtain the elemental composition by Particle Induced X-ray Emission
155 (PIXE) technique, using a properly collimated proton beam and scanning the deposits in steps corresponding to 1-hour
156 aerosol deposit (details in Calzolari et al., 2015). In this work, fine and coarse elemental concentrations determined by
157 PIXE analysis were added up to obtain PM10 concentrations with hourly resolution as low time resolution PM10 samples
158 were also available.

159

160 *2.3 Aerosol light-absorption coefficient measurements*

161 The aerosol absorption coefficient (b_{ap}) at the 4 wavelengths $\lambda = 405, 532, 635$ and 780 nm was measured on both low
162 and high time resolution samples with the home-made polar photometer PP_UniMI (Vecchi et al., 2014; Bernardoni et
163 al., 2017c). Low time resolution optical measurements were performed on PTFE filters since their physical characteristics
164 can be considered more similar to polycarbonate filters used by the streaker sampler. Indeed, Vecchi et al. (2014) reported
165 that aerosol absorption measurements on samples collected in parallel on quartz-fibre and PTFE filters showed significant
166 differences, which were mainly ascribed to sampling artefacts due to organics affecting quartz-fibre filters. For high time
167 resolution samples, b_{ap} was measured only in the fine fraction collected on polycarbonate filters, since absorption from
168 the Kapton foil on which the coarse fraction was collected did not allow b_{ap} assessment. Anyway, b_{ap} values in PM2.5
169 and PM10 were expected to be fairly comparable, as most of the contribution to aerosol absorption in atmosphere is
170 typically given by particles in the fine fraction at heavily polluted urban sites like Milan. To verify this assumption, high
171 time resolution b_{ap} data in PM2.5 were averaged on the time scale of low time resolution b_{ap} in PM10 for comparison.
172 They turned out to be in good agreement, between 11 % and 13 % depending on the λ , except for b_{ap} at $\lambda=405$ nm that
173 showed a higher difference (27 %) but with most data (83 %) within experimental uncertainties. To take into account for
174 this difference, b_{ap} data at $\lambda=405$ nm were homogenised before their insertion into the model, following the criterion used
175 for chemical species (for further detail about homogenisation procedure, see Sect. 2.4 and Sect. 2.5).

176 Uncertainties on b_{ap} were about 15 %. MDL was in the range 1-10 Mm⁻¹ depending on sampling duration and wavelength.
177 Optical system stability was checked during the measurement session, evaluating the reproducibility of the measure of a
178 blank test filter. Laser stability was also checked at least twice a day and the recorded intensities were used to normalise
179 blank and sampled filters analysis.

180

181 *2.4 Model description*



182 Multivariate receptor models (Henry, 1997) represent one of the most widespread and robust approaches used to carry
183 out source apportionment studies for atmospheric aerosol (Belis et al., 2014 and 2015). In particular, the Positive Matrix
184 Factorization (Paatero and Tapper, 1994; Paatero, 1997) has been extensively used in the literature and, afterwards, the
185 Multilinear Engine ME2 (Paatero, 1999 and 2000) introduced the possibility of solving all kinds of multilinear and quasi-
186 multilinear problems. The fundamental principle is the mass conservation between the emission source and the receptor
187 site: focusing on the aerosol properties at the receptor site (i.e. measuring a large number of chemical constituents in
188 different samples), a mass balance analysis can be performed to identify the factors influencing aerosol concentrations
189 (Hopke, 2016). Factors can be subsequently interpreted as the main sources impacting the site, exploiting the
190 characterisation of the most relevant sources in the investigated area or the adoption of fingerprints available from
191 previous literature works (Belis et al., 2014). Referring to the input data as matrix X (matrix elements x_{ij}), the chemical
192 profile of the factors as matrix F (matrix elements f_{kj}), and the time contribution of the factors as matrix G (matrix elements
193 g_{ik}), the main equation of a bilinear problem can be written as follows:

$$194 \quad x_{ij} = \sum_{k=1}^P g_{ik} f_{kj} + e_{ij} \quad (1)$$

195 where the indices i , j , and k indicate the sample, the species, and the factor, respectively; P is the number of factors and
196 the matrix E (matrix elements e_{ij}) is composed by the residuals, i.e. the difference between measured and modelled values.
197 In this way, a system of $N \times M$ equations is established, where N is the number of samples and M is the number of species.
198 The solution of the problem is computed minimising the object function Q defined as:

$$199 \quad Q = \sum_{i=1}^N \sum_{j=1}^M \left(\frac{e_{ij}}{\sigma_{ij}} \right)^2 \quad (2)$$

200 where σ_{ij} are the uncertainties related to the input data.

201 The multi-time receptor model was developed in order to use each data value in its original time schedule, without
202 averaging the high time resolution data or interpolating the low time resolution data (Zhou et al., 2004; Ogulei et al.,
203 2005). The main Eq. (1) is consequently modified as below:

$$204 \quad x_{sj} = \frac{1}{t_{s2} - t_{s1} + 1} \sum_{k=1}^P f_{kj} \sum_{i=t_{s1}}^{t_{s2}} g_{ik} \eta_{jm} + e_{sj} \quad (3)$$

205 where the indices s , j , and k indicate the sample, the species and the factor respectively; P is the number of factors; t_{s1} and
206 t_{s2} are the starting and ending time for the s -th sample in time units (i.e. the shortest sampling interval, that is 1 hour for
207 the dataset of this work) and i represents one of the time units of the s -th sample. η_{jm} are adjustment factors for chemical
208 species replicated with different time resolution and measured with different analytical methods (represented by the
209 subscript m).



210 If η is close to unity, species concentration measured by different analytical approaches can be considered in good
211 agreement; non-replicated species have adjustment factors set to unity by default. In this work, the adjustment factors
212 were always set to unity in the model; to take into account the use of different aerosol samplers (i.e. low volume sampler
213 with EPA inlet and streaker sampler) and different analytical techniques to obtain the elemental composition (i.e. ED-
214 XRF and PIXE), concentrations of replicated species with different time resolution were homogenised before inserting
215 them into the input matrix X, as will be explained in Sect. 2.5. Applying this data treatment procedure, it is possible to
216 avoid the check of the consistency, from the experimental point of view, of the η values calculated by the model.
217 Otherwise, this step should always be performed after running the model.

218 In the multi-time model a regularisation equation is introduced, since some sources could contain few or no species
219 measured with high time resolution:

$$220 \quad g_{(i+1)k} - g_{ik} = 0 + \varepsilon_i \quad (4)$$

221 where ε_i represent the residuals.

222 As already pointed out in a previous work (Ogulei et al., 2005), a weighting parameter for low resolution species (24 or
223 12 hours in this work) might be necessary; in this study, it was implemented in the equations and set at 0.5 for strong
224 species (not applied to weaker species as Na, Mg, and Cr, see Sect. 2.5).

225 Equations (3) and (4) were solved using the Multilinear Engine (ME) program (Paatero, 1999). In Eq. (2), the object
226 function Q takes into account residuals from the main Eq. (3) and from the auxiliary equations (regularisation Eq. (4),
227 normalisation equation, pulling equations, and constraints).

228 In this work, the multi-time model implemented by Crespi et al. (2016) is used. Using this model implementation,
229 constraints can be inserted in the model and the bootstrap analysis is also performed to evaluate the robustness of the final
230 solution.

231

232 *2.5 Input data*

233 As already mentioned in Sect. 2.4, instead of using adjustment factors in the model (all set equal to one), concentrations
234 of replicated species with different time resolution were pre-homogenised and then inserted into the input matrix X.
235 Concentration data with longer sampling interval (24 and 12 hours in this work) were considered as benchmark, since
236 analytical techniques usually show a better accuracy on concentration values far from MDL (i.e. samples collected on
237 longer time intervals) (Zhou et al., 2004; Ogulei et al., 2005).

238 Variables were then classified as weak and strong based on the signal to noise ratio (S/N) criterion (Paatero, 2015). For
239 hourly data only strong variables ($S/N \geq 1.2$) were considered; for low time resolution data also weaker variables as Na,



240 Mg and Cr (with S/N equal to about 0.8), that resulted strong variables in hourly samples, were also inserted (with
241 uncertainties comparable to concentration values) in order to avoid the exclusion of too many data. Indeed, excluding
242 these low time resolution variables from the analysis gave rise to artificial high values in the source time contribution
243 matrix for those sources traced by these species (in this case it was particularly important for aged sea salt traced by Na
244 and Mg, see Sect. 3.1); this oddity was already reported by Zhou et al. (2004).

245 Variables with more than 20 % of the concentration data below MDL values were omitted from the analysis (Ogulei et
246 al., 2005). The procedure described in Polissar et al. (1998) was followed to treat uncertainties and below MDL data. In
247 general, missing concentration values were estimated by linear interpolation of the measured data and their uncertainties
248 were assumed as three times this estimated value (Zhou et al., 2004; Ogulei et al., 2005). As for summertime levoglucosan
249 data (not available), the approach was to include them as below MDL data and not as missing data following Zhou et al.
250 (2004), who underlined that the multi-time model is more sensitive to missing values than the original PMF model. In
251 order to avoid double counting, in this study S was chosen as input variable instead of SO_4^{2-} as it was determined on both
252 low time- and high time-resolution samples (by XRF and PIXE analysis, respectively, see Calzolari et al., 2008). However,
253 elemental SO_4^{2-} and S concentrations showed a high correlation (correlation coefficient $R=0.98$) and the Deming
254 regression gave a slope of 2.69 ± 0.13 (sulphate vs. sulphur) with an intercept of -198 ± 82 , i.e. compatible with zero
255 within 3 standard deviations. The slight difference (of the order of 10%) between the estimated slope and the SO_4^{2-} -to-S
256 stoichiometric coefficient (i.e. 3) can be ascribed to either a small fraction of insoluble sulphate or to the use of different
257 analytical techniques.

258 PM10 mass concentrations were included in the model with uncertainties set at four times their values (Kim et al., 2003).
259 In the end, 22 low time resolution variables (PM10 mass, Na, Mg, Al, Si, S, K, Ca, Cr, Mn, Fe, Cu, Zn, Pb, EC, OC,
260 levoglucosan, NO_3^- , $b_{\text{ap}} 405\text{nm}$, $b_{\text{ap}} 532\text{nm}$, $b_{\text{ap}} 635\text{nm}$, $b_{\text{ap}} 780\text{nm}$) and 17 hourly variables (Na, Mg, Al, Si, S, K, Ca, Cr,
261 Mn, Fe, Cu, Zn, Pb, $b_{\text{ap}} 405\text{nm}$, $b_{\text{ap}} 532\text{nm}$, $b_{\text{ap}} 635\text{nm}$, $b_{\text{ap}} 780\text{nm}$) were considered.

262 Finally, the input matrix consisted in 386 samples and the total number of time units was 1117. The analysis was
263 performed in the robust mode; lower limit for G contribution was set to -0.2 (Brown et al., 2015) and the error model
264 $\text{em}=-14$ was used for the main equation with C_1 = input error, $C_2=0.0$ and $C_3=0.1$ (Paatero, 2012).

265 To the authors' knowledge, this is the first time that b_{ap} at different wavelengths has been introduced in the multi-time
266 model and used to more robustly identify the sources; moreover, the optical information has been also exploited to retrieve
267 additional information such as the Absorption Ångström Exponent (α) of the sources and MAC values in an original way.

268

269 3. Results and discussion

270 3.1 Source apportionment with multi-time model

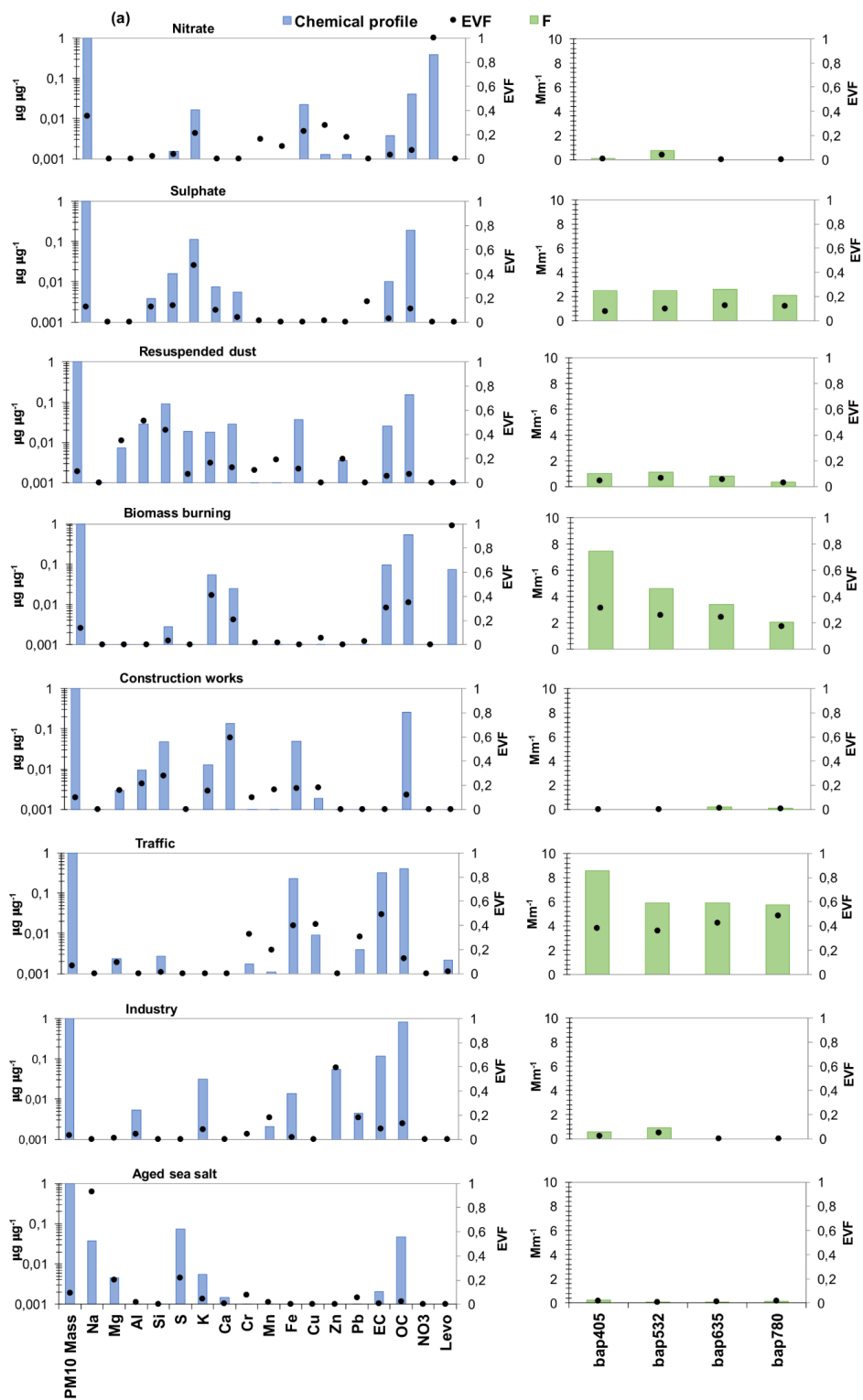


271 Different number of factors (5-10) were explored; after 30 convergent runs, an 8-factor base-case solution corresponding
272 to the lowest Q value (2086.88) was firstly selected (see Fig. S1 in the Supplement). It is important to notice that the
273 model was run using all variables (chemical + optical) as explained in Sect. 2.5. A lower or higher number of factors
274 caused ambiguous chemical profiles as the physical interpretation highlighted clearly mixed sources for a lower number
275 of factors or unique factors in case of more factors (i.e. Pb for 9 factors); moreover, inconsistent mass closure was detected
276 increasing the number of factors (e.g. the sum of species contribution was up to 25 % higher than the mass for the 10-
277 factor solution). In the 8-factor base case solution, the mass was well reconstructed by the model ($R^2=0.98$), with a slope
278 of 0.98 ± 0.02 and negligible intercept= $0.51 \pm 0.89 \mu\text{g m}^{-3}$.

279 The factor-to-source assignment process was based on both the Explained Variation (EV) values - which are typically
280 higher for chemical tracers (Lee et al., 1999; Paatero, 2010) - and the physical consistence of factor chemical profiles. In
281 chosen solution, the not explained variation was lower than 0.25 for all variables. The scaled residuals showed a random
282 distribution of negative and positive values in the ± 3 range.

283 Using EV and chemical profiles reported in Fig. S1 (a), the 8 factors were tentatively assigned to specific atmospheric
284 aerosol sources: nitrate, sulphate, resuspended dust, biomass burning, construction works, traffic, industry, and aged sea
285 salt. In Table S1 (in the Supplement) absolute and relative average source contributions to PM10 mass are reported.

286 Although the above mentioned base-case solution was a satisfactory representation of the main sources active in the area
287 (as reported in previous works, see e.g. Marcazzan et al., 2003; Vecchi et al., 2009 and 2018; Bernardoni et al., 2011 and
288 2017a; Amato et al., 2016), the chemical profiles of some factors could be improved exploring rotated solutions. The
289 most relevant case was represented by aged sea-salt where typical diagnostic ratios such as Mg/Na and Ca/Na were not
290 well reproduced (in bulk sea water equal to 0.12 and 0.04, respectively, as reported e.g. in Seinfeld and Pandis, 2006) and
291 the chemical profile itself was too much impacted by the presence of Fe compared to bulk sea water composition.
292 Therefore, the above-mentioned diagnostic ratios were here used as constraints and Fe was maximally pulled down in the
293 chemical profile. The effective increase in Q was of about 61 units ($Q=2147$), with a percentage increase of about 3 %;
294 as a rule of thumb, an increase in the Q value of a few tens is generally acceptable (Paatero and Hopke, 2009). It is
295 noteworthy that an improvement in the chemical profiles was achieved with negligible differences— compared to the base-
296 case solution – as for all other relevant features of the solution (i.e. EV, residuals, mass reconstruction, source
297 apportionment). Therefore, an 8-factor constrained solution was considered the most physically reliable; results are
298 presented in Table 1 and Fig. 1 and discussed in detail in the following.





300 Figure 1: (a) Chemical profiles of the 8-factor constrained solution (b) b_{ap} apportionment of the 8-factor constrained
 301 solution.
 302

303

Factors	Summer [$\mu\text{g m}^{-3}$]	Winter [$\mu\text{g m}^{-3}$]	Total [$\mu\text{g m}^{-3}$]
Nitrate	3.6 (15 %)	21.1 (44 %)	10.2 (31 %)
Sulphate	6.3 (26 %)	8.1 (17 %)	7.0 (21 %)
Re-suspended dust	4.6 (19 %)	1.7 (4 %)	3.5 (11 %)
Biomass burning	0.32 (1 %)	8.3 (17 %)	3.3 (10 %)
Construction works	5.9 (24 %)	3.4 (7 %)	4.9 (15 %)
Traffic	1.4 (6 %)	2.2 (5 %)	1.7 (5 %)
Industry	0.86 (4 %)	1.2 (3 %)	1.0 (3 %)
Aged sea salt	1.4 (6 %)	1.8 (4 %)	1.6 (5 %)

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Table 1: Absolute and relative average source contributions to PM10 mass in the 8-factor constrained solution.

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In the factor interpreted as nitrate the explained variation is fully ascribed to NO_3^- . This factor contains a significant fraction of nitrate in the chemical profile (39 %) and all nitrate is accounted for only in this factor. This source is by large the most significant source at the investigated site, accounting for about 31 % over the whole campaign (a similar estimate – 26 % - was reported by Amato et al. (2016) during the AIRUSE campaign in Milan in 2013) raising up to 44 % during wintertime (comparable to 37 % reported by Vecchi et al. (2018)). Indeed, the Po valley is well-known for experiencing very high nitrate concentrations during wintertime (Vecchi et al., 2018; and references therein) because of large emissions of gaseous precursors related to urban and industrial activities, wood burning used for residential heating, high ammonia levels due to agricultural fields manure and – last but not the least – poor atmospheric dispersion conditions.

The factor associated to sulphate shows $\text{EV}=0.47$ for S, which contributes to the chemical profile for about 11 % (corresponding to 33 % as SO_4^{2-} and 45 % as ammonium sulphate). Due to the secondary origin of the aerosol associated to this factor, it is not surprising to find also a significant OC contribution (19 %); indeed, as already pointed out for the same location by Vecchi et al. (2018), in Milan aerosol is impacted by highly oxygenated components due to aging processes favoured by strong atmospheric stability (Vecchi et al., 2019). This factor accounts for 21 % of the PM10 mass. It is noteworthy that, on average, factors associated to secondary compounds (i.e. nitrate and sulphate) together account for 52 % of the PM10 mass in the Milan atmosphere for the period under investigation.

The factor identified as re-suspended dust is mainly characterised by high EVs and contributions coming from Al, Si and Mg, i.e. crustal elements. The Al/Si ratio is 0.31, very similar to the literature value for average crust composition (Mason, 1966); the relatively high contribution of OC in the chemical profile (15 %) and the presence of EC (about 2.6 %), indicate that there is very likely a mixing with road dust (Thorpe and Harrison, 2008). This source accounts for about 11 % of the PM10 mass.



327 The factor identified as biomass burning is characterised by high EV for levoglucosan (0.98), a known tracer for this
328 source as it is generated by cellulose pyrolysis; EV higher than 0.3 are also found for K, OC, and EC. In the source
329 chemical profile, OC contributes for 54 %, EC for 10 %, levoglucosan for 7 %, and K for 5 %. The average biomass
330 burning contribution during this campaign is 10 % (up to 17 % in wintertime). Anticipating the discussion presented in
331 detail in Sect. 3.2, it is worth noticing that in this factor there is also the second largest contribution to the aerosol
332 absorption coefficient after traffic.

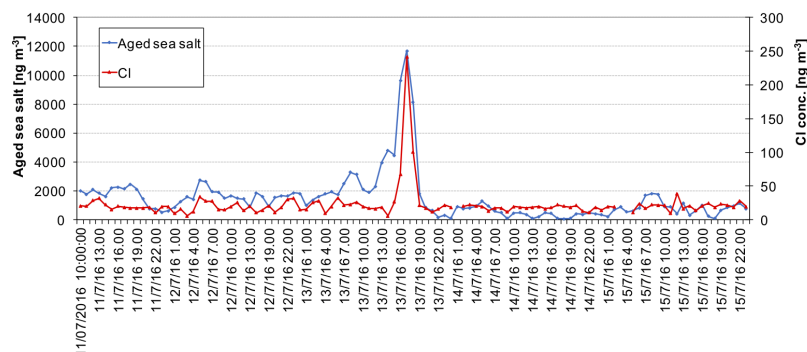
333 The factor presenting a high (0.60) EV for Ca is associated to construction works, following literature works (e.g. Vecchi
334 et al., 2009; Bernardoni et al., 2011; Dall'Osto, 2013; Crilley et al., 2017; Bernardoni et al., 2017a; and references therein).
335 Major contributors to the chemical profile are Ca (13 %), OC (26 %), Fe and Si (5 % each). This factor accounts on
336 average for 15 % to PM₁₀. As already mentioned, during the campaign a not negligible contribution from this source
337 could be expected, due to the presence of a construction building site nearby the monitoring location.

338 In the factor here assigned to traffic (primary contribution), EV larger than 0.3 are found for EC, Cu, Fe, Cr, and Pb. The
339 highest relative contributions in terms of mass in the chemical profile are given by OC (41 %), EC (32 %), Fe (23 %),
340 and Cu (1 %). The OC-to-EC ratio is about 1.3 which is consistent with a primary traffic contribution (Giugliano et al.,
341 2005; Bernardoni et al., 2011). This traffic (primary) contribution over the whole dataset accounts for 5 % of the PM₁₀
342 mass with a slightly lower absolute contribution in summer (see Table 1). This contribution is comparable to the
343 percentage (7 %) reported by Amato et al. (2016) for exhaust traffic emissions but it is lower than our previous estimates,
344 i.e. 15 % in 2006 in PM₁₀ and 12 % in PM₁ recorded in winter 2012, as reported in Bernardoni et al. (2011) and Vecchi
345 et al. (2018), respectively. However, the current estimate seems to be still reasonable when considering the efforts done
346 in latest years to reduce vehicles exhaust particle emissions and the fraction of secondary nitrate to be added to account
347 for the overall traffic impact; indeed, a significant traffic contribution due to nitrate should be accounted for the relevant
348 nitrogen oxides and ammonia emissions from agriculture in the region (INEMAR ARPA-Lombardia, 2018).
349 Unfortunately, the non-linearity of the emission-to-ambient concentration levels relationship and the high uncertainties
350 in emission inventories still prevent a robust estimate of this secondary contribution to total traffic exhaust emissions. In
351 Sect. 3.2, it will be shown that traffic is the largest contributor to aerosol absorption coefficient, a result that reinforces
352 the interpretation of this factor as a traffic emission source.

353 The industry factor shows high EV for Zn (0.59); the second highest EV is related to Mn (0.13). Previous studies at the
354 same sampling site identified these elements as tracers for industrial emissions (e.g. Vecchi et al., 2018; and references
355 therein). The chemical profile is enriched by heavy metals and, after traffic, it is the profile with the highest share of Cr,
356 Mn, Fe, Cu, Zn, and Pb (explaining about 8 % of the total PM₁₀ mass in the profile). The industry contribution is not
357 very high in the urban area of Milan, accounting for 3 % on average.

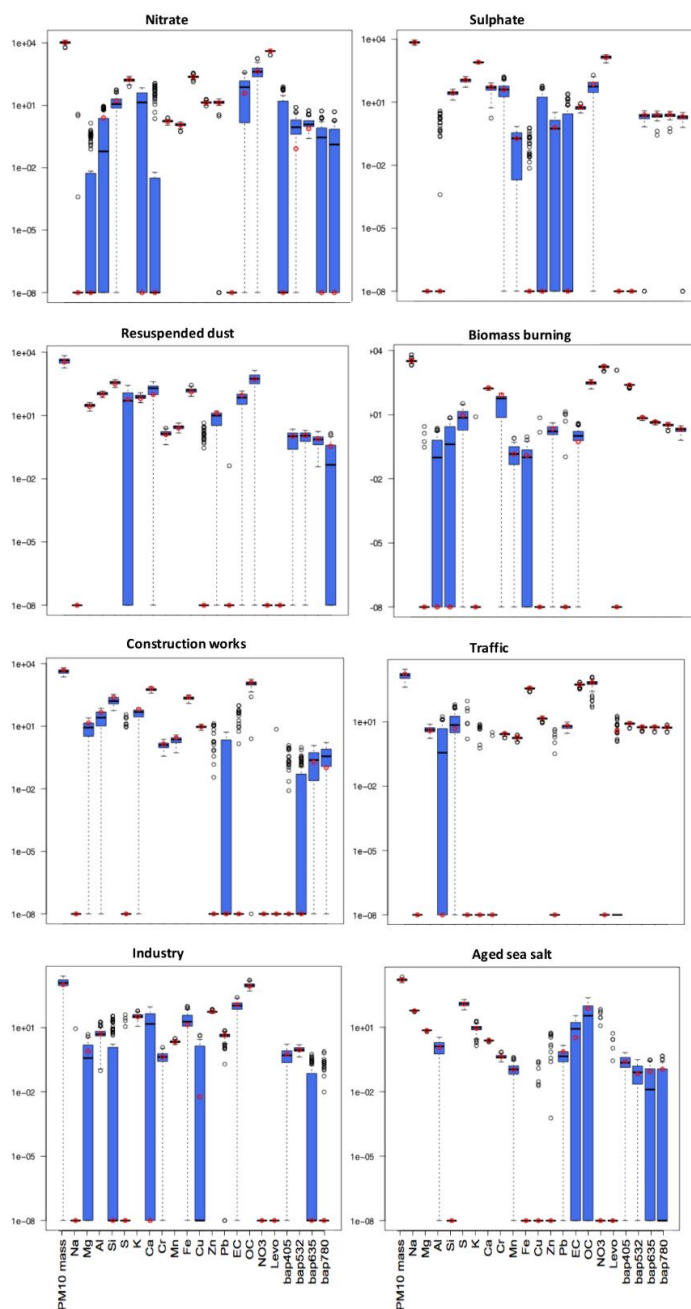


358 The factor interpreted as aged sea salt has high EV of Na (0.93) and this element is - as a matter of fact - present only in
359 this factor chemical profile. To check the physical consistency of this assignment and considering that Milan is about 120
360 km away from the nearest sea coast, back-trajectories frequencies were calculated through the NOAA HYSPLIT
361 trajectory model (Draxler and Hess, 1998; Stein et al., 2015; Rolph et al., 2017). Temporal patterns of Cl concentrations
362 (not inserted in the multi-time analysis as being a weak variable) during episodes were exploited to further confirm the
363 factor-to-source association. As an example, a very short event (13/07 h. 16-18) highlighted by the model and representing
364 the highest sea salt contribution during summer is here analysed in further detail. Before and during the sea salt event, air
365 masses showed a clear origin from south-west compatible with Ligurian sea while soon after the event, there was a rapid
366 change of wind direction (Fig. S2, in the Supplement). These hours were characterised by an average high wind speed of
367 $4.8 \pm 1.7 \text{ m s}^{-1}$ (with a maximum peak of 9.5 m s^{-1}) compared to $1.9 \pm 1.0 \text{ m s}^{-1}$ average wind speed recorded during the
368 summer campaign. In addition, Cl concentration and aged sea salt pattern show an evident temporal coincidence in peak
369 occurrence during the event (Fig. 2), thus confirming the correctness of the source identification. During this episode,
370 only the Cl coarse fraction increased (Fig. S3, in the Supplement) and reached about 90 % of total PM10 Cl concentration;
371 Cl/Na ratio was 0.38 ± 0.05 , consistent with an aging of marine air masses during advection showing the typical Cl
372 depletion due to the interaction between sea salt particles and polluted air masses (Seinfeld and Pandis, 2006).



373
374 Figure 2: Temporal patterns of aged sea salt source retrieved from the multi-time model and Cl concentrations measured
375 in atmosphere.

376
377 Bootstrap analysis was performed to evaluate the uncertainties associated to source profiles (Crespi et al., 2016). 100 runs
378 were carried out (see Fig. 3, values expressed in ng m^{-3} or Mm^{-1} on a logarithmic scale); factors were well mapped, with
379 Pearson coefficient always higher than 0.97, and tracers for each source showed small interquartile range, supporting the
380 goodness of the solution presented in this work.



381

382 Figure 3: Box plot of the bootstrap analysis on the 8-factor constrained solution. The red dots represent the output values

383 of the solution of the model; the black lines the medians from the bootstrap analysis; the blue bars the 25th and 75th

384 percentile; the dotted lines the interval equal to 1.5 the interquartile range and the black dots the outliers from this interval.



385

386 *3.2 Improving source apportionment with optical tracers*

387 First of all, the use of the absorption coefficient determined at different wavelengths as input variable in the multi-time
388 model, strengthened the identification of the sources, suggesting that it can be exploited when specific chemical tracers
389 are not available (e.g. levoglucosan for wood burning). To prove that, a separate source apportionment study was
390 performed with EPA PMF 5.0 (Norris et al., 2014), introducing only hourly elemental concentrations from samples
391 collected by the streaker sampler and hourly b_{ap} at different λ measured by PP_UniMI on the same filters. Streaker
392 samples typically lack of a complete chemical characterisation; in particular, important chemical tracers such as
393 levoglucosan and EC are not available. In this analysis, b_{ap} assessed at different wavelengths resulted particularly useful
394 for the identification of the biomass burning factor that explained a significant percentage of the b_{ap} itself (from 25 % to
395 35 % depending on λ) (Fig. S4, in the Supplement); without this additional information, the factor-to-source assignment
396 would be otherwise based only on the presence of elemental potassium while being well-known that potassium cannot be
397 considered an unambiguous tracer as it is emitted by a variety of sources (see for example Pachon et al., 2013; and
398 references therein).

399 As for the multi-time model, results showed that the absorption coefficient contribution was higher than 45 % in the factor
400 labelled as traffic, highlighting the importance of exhaust emissions in a factor that would be otherwise characterised
401 mainly on elements related to non-exhaust emissions (Cu, Fe, Cr).

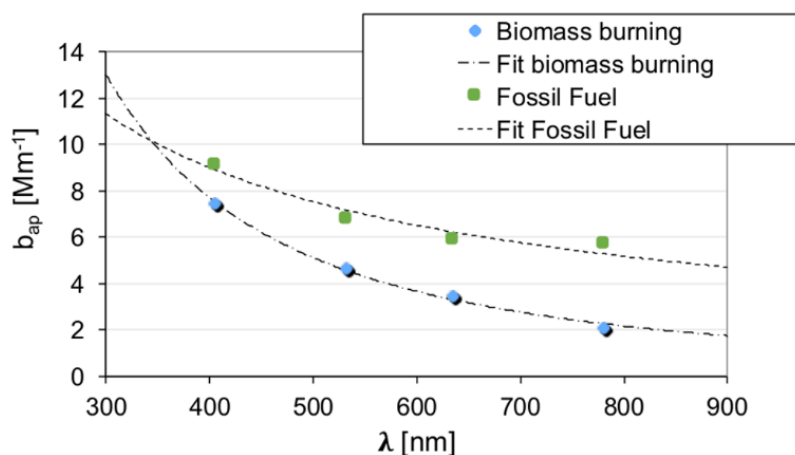
402 Furthermore, the two factors identified as biomass burning and traffic are the main contributors to aerosol absorption in
403 atmosphere and show significant EV values. Traffic accounts for 55 % of b_{ap} at 780 nm and 42 % at 405 nm; biomass
404 burning accounts for 20 % and 36 % at 780 and 405 nm, respectively. The Explained Variation (EV) of b_{ap} has the
405 maximum value at 405 nm for biomass burning (0.32) and at 780 nm for traffic (0.49), showing the tendency to decrease
406 and increase with the wavelength, respectively. The other six sources are less relevant in terms of EV values and overall
407 contribute for less than 30%.

408 It is noteworthy that opposite to the approach used in source apportionment optical models, like the widespread
409 Aethalometer model (Sandradewi et al., 2008) and MWAA model (Massabò et al., 2015; Bernardoni et al., 2017b), no a-
410 priori information about the Ångström Absorption Exponent (α) of the fossil fuel and biomass burning sources was
411 introduced in the multi-time model; instead, an estimate for its value can be directly retrieved from the obtained solution.

412 It has to be mentioned that optical models are based on a two-source hypothesis (i.e. biomass burning and fossil fuel
413 emissions). Hereafter, in order to compare multi-time model and optical models results, contributions due to traffic and
414 industry (i.e. emissions most likely connected to fossil fuel usage) were added up and labelled as “fossil fuel emissions”.



415 In Fig. 4 the wavelength dependence of b_{ap} for the biomass burning and the fossil fuel profiles obtained with the multi-
416 time model is shown; the fitting procedure considering $b_{ap} \propto \lambda^{-\alpha}$, gives α_{BB} (α biomass burning) = 1.83 and α_{FF} (α fossil
417 fuels) = 0.80. The bootstrap analysis allowed to estimate the range of variability of α values, considering the 25th and 75th
418 percentile: 0.78-0.88 for α_{FF} and 1.65-1.88 for α_{BB} .



419

420 Figure 4: b_{ap} dependence on λ for biomass burning and fossil fuel emissions.

421

422 The value of α_{BB} obtained in this work is very similar to 1.86 found for biomass burning by Sandradewi et al. (2008) and
423 1.8 obtained by Massabò et al. (2015) who used also independent ¹⁴C measurements for checking.

424 The α_{FF} value (assumed to be equal to α_{BC} in source apportionment optical models) obtained in this work is in the range
425 0.8-1.1 typically reported in optical source apportionment studies (e.g. Bernardoni et al., 2017b; and references therein).

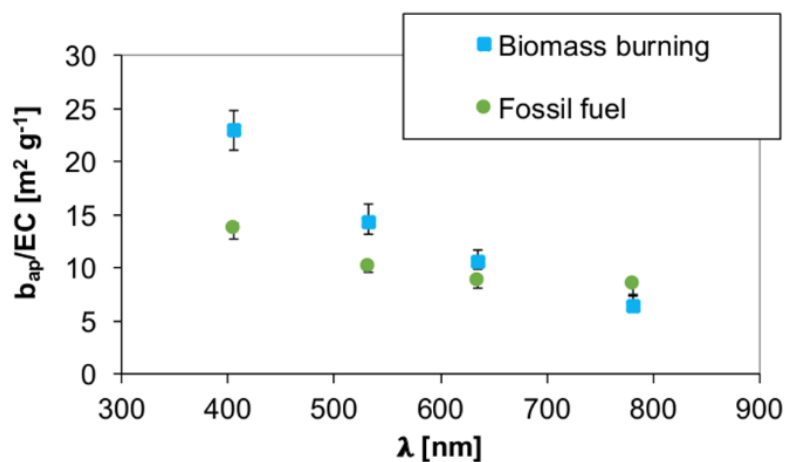
426 It is also consistent with the atmospheric α value obtained during the summer campaign, when biomass burning was
427 negligible (impacting 1 % of the total PM10 mass from the multi-time source apportionment). The assessment of α_{BC} is
428 still an issue and both experimental and simulation studies are in progress to reduce uncertainties and give a better
429 evaluation of this relevant optical parameter; literature works (e.g. Gyawali et al., 2009; Fischer and Smith, 2018; Liu et
430 al. 2018) report α_{BC} values ranging from 0.6 to 1.4 for both coated and bulk particles with a strong dependence from
431 particle size distribution.

432 Results here reported allow also to study the relationship between the absorption coefficient and the mass of black carbon,
433 i.e. the so called Mass Absorption Cross section (MAC) at different wavelengths. The $MAC(\lambda) = b_{ap}(\lambda)/BC$ relationship
434 assumes that black carbon (BC) is the only light-absorbing species present; however, this assumption is not always valid,
435 since mineral dust and brown carbon (BrC) can significantly contribute to aerosol absorption. During our monitoring



436 campaign, no significant contribution from mineral dust was observed; opposite, biomass burning was proved to be a
437 relevant source so that BrC was certainly a significant contributor (Fuzzi et al., 2015) as also suggested by $\alpha_{BB} = 1.83$ in
438 the biomass burning factor. The possible overestimation of BC when total b_{ap} is ascribed to BC only is usually minimised
439 choosing a wavelength higher than 600 nm, exploiting the spectral dependence of absorption from different aerosol
440 compounds (Petzold et al., 2013).

441 EC concentration retrieved from the chemical profiles (see Fig. 1) was used as a proxy for BC to estimate source-
442 dependent $b_{ap}(\lambda)$ -to-BC ratio. Results are represented in Fig. 5. It is noteworthy that here this ratio is intentionally not
443 indicated as MAC, since overestimation of the BC absorption especially at lower λ might occur (see previous discussion).
444 BrC is expected to give a small contribution in the fossil fuel source; therefore, the best approximation for $MAC(\lambda)$ values
445 are likely the $b_{ap}(\lambda)$ -to-BC ratios observed in the fossil fuel source at our monitoring site. They resulted to be $13.7 \text{ m}^2 \text{ g}^{-1}$
446 at $\lambda = 405 \text{ nm}$; $10.2 \text{ m}^2 \text{ g}^{-1}$ at $\lambda = 532 \text{ nm}$; $8.8 \text{ m}^2 \text{ g}^{-1}$ at $\lambda = 635 \text{ nm}$; $8.6 \text{ m}^2 \text{ g}^{-1}$ at $\lambda = 780 \text{ nm}$. At $\lambda = 550 \text{ nm}$ Bond and
447 Bergstrom (2006) report $MAC = 7.5 \pm 1.2 \text{ m}^2 \text{ g}^{-1}$ for uncoated fresh emitted particles and MAC values in polluted regions
448 ranging from 9 to $12 \text{ m}^2 \text{ g}^{-1}$, attributable to absorption enhancement due to particles coating. The MAC estimate obtained
449 in this work from multi-time model at 532 nm is comparable to literature values and it confirms the importance of aging
450 processes in atmosphere on the optical properties of particles.



451

452 Figure 5: b_{ap} -to-EC ratio dependence on λ for biomass burning and fossil fuel emissions. Error bars represent the 25th and
453 75th percentile retrieved from the bootstrap analysis.

454

455 Ratios in Figure 5 are clearly less comparable at $\lambda=405 \text{ nm}$ (see also table 2S); this result is explained by the significant
456 contribution of BrC to b_{ap} at this wavelength in the biomass burning factor.



457 No seasonal differences in the atmospheric ratios were observed but at $\lambda = 405$ nm (see table S2, in the Supplement), for
458 which winter values are higher than summer ones (17.8 ± 0.4 and 14.2 ± 0.5 , respectively); this result can be explained
459 considering the influence of biomass burning emissions on BrC concentration in atmosphere during the winter season.
460 From the outputs of the modelling approach here proposed, the apportionment of the biomass burning and fossil fuel
461 contributions to b_{ap} at different wavelengths was also obtained. As expected, the relative contribution to the total
462 reconstructed b_{ap} ascribed to the biomass burning factor decreases with increasing λ , opposite to the contribution from
463 fossil fuel combustion which gives the highest contribution at 780 nm (table 1); in addition, the latter contribution prevails
464 at all wavelengths at the investigated site.
465

	$\lambda = 405$ nm	$\lambda = 532$ nm	$\lambda = 635$ nm	$\lambda = 780$ nm
Biomass burning	36 % (31 %-36 %)	29 % (25 %-30 %)	26 % (23 %-27 %)	20 % (16 %-22 %)
Fossil fuels	45 % (41 %-46 %)	43 % (39 %-44 %)	45 % (41 %-47 %)	55 % (48 %-55 %)

466 Table 2: Average contribution to total reconstructed b_{ap} for the biomass burning and fossil fuel factors; in parenthesis 25th
467 and 75th percentile are reported.

468

469 4. Conclusions

470 The multi-time model implemented through Multilinear Engine (ME2) script allowed the analysis of experimental data
471 collected at different time scales, coupling the detailed chemical speciation at low time resolution and the temporal
472 information given by high time resolution samples. The effect of the introduction of the aerosol absorption coefficient
473 (b_{ap}) measured at different wavelengths in the modelling process was investigated and gave promising results. First of all,
474 a more robust identification of sources was provided; secondly, it paved the way to the retrieval of optical apportionment
475 and optical characterisation of the sources (e.g. estimate of source-specific Ångström Absorption Exponent - α - and
476 MAC at different wavelengths). It is worthy to note that – at the state of the art – in source apportionment optical models
477 (e.g. Aethalometer model) values for α related to fossil fuel emissions and wood burning are fixed by the modeller thus
478 carrying a large part of the uncertainties in the model results. The original approach described in this work can be applied
479 to any source apportionment study using any suitable dataset (not necessarily with multi-time resolution). Besides the
480 traditional source apportionment, the impact of different sources on the aerosol absorption coefficient was estimated; this
481 piece of information can be very useful to formulate strategies of pollutants abatement, in order to improve air quality
482 and to face climate challenges. In particular, at the investigated site secondary compounds constituted the highest
483 contribution in terms of PM10 mass (52 % on average), while the two factors identified as biomass burning and traffic



484 were found to be the most significant contributors to aerosol absorption in atmosphere, in agreement with available
485 literature works.

486

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493

494 **Data availability.**

495 The data in the study are available from the authors upon request (roberta.vecchi@unimi.it).

496

497 **Supplement.**

498 The supplement related to this article is available online

499

500 **Author contributions.**

501 ACF performed streaker sampling and related optical analysis, implemented the advanced model, analysed the results,
502 and drafted the paper. GV contributed to model implementation and data reduction. VB, SV, and REP carried out the
503 sampling campaign on filters, performed the optical measurements and data analysis. GC, SN, and FL performed PIXE
504 analysis and data reduction. DM and PP carried out ionic characterisation on filters and data analysis. RV was responsible
505 for the design and coordination of the study, the synthesis of the results and the final version of the paper. All authors
506 contributed to the interpretation of the results obtained with the new approach here described and revised the manuscript
507 content giving a final approval of the version to be submitted.
508

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510 **Competing interests.**

511 The authors declare that they have no conflict of interest.

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743 **List of Captions**

744 Figure 1: (a) Chemical profiles of the 8-factor constrained solution (b) b_{ap} apportionment of the 8-factor constrained
745 solution.

746 Figure 2: Temporal patterns of aged sea salt source retrieved from the multi-time model and Cl concentrations measured
747 in atmosphere.

748 Figure 3: Box plot of the bootstrap analysis on the 8-factor constrained solution. The red dots represent the output values
749 of the solution of the model; the black lines the medians from the bootstrap analysis; the blue bars the 25th and 75th
750 percentile; the dotted lines the interval equal to 1.5 the interquartile range and the black dots the outliers from this interval.

751 Figure 4: b_{ap} dependence on λ for biomass burning and fossil fuel emissions.

752 Figure 5: b_{ap}-to-EC ratios dependence on λ for biomass burning and fossil fuel emissions. Error bars represent the 25th
753 and 75th percentile retrieved from the bootstrap analysis.

754

755 Table 1: Absolute and relative average source contributions to PM₁₀ mass in the 8-factor constrained solution.

756 Table 2: Average contribution to total reconstructed b_{ap} for the biomass burning and fossil fuel factors; in parenthesis
757 25th and 75th percentile are reported.

758