

1 Exploiting multi-wavelength aerosol absorption coefficients in a 2 multi-time resolution source apportionment study to retrieve 3 source-dependent absorption parameters

4 Alice C. Forello¹, Vera Bernardoni¹, Giulia Calzolai², Franco Lucarelli², Dario Massabò³, Silvia
5 Nava², Rosaria E. Pileci^{1,a}, Paolo Prati³, Sara Valentini¹, Gianluigi Valli¹, Roberta Vecchi^{1,*}

6 ¹Department of Physics, Università degli Studi di Milano and National Institute of Nuclear Physics INFN-Milan, via
7 Celoria 16, Milan, 20133, Italy

8 ²Department of Physics and Astronomy, Università di Firenze and National Institute of Nuclear Physics INFN-Florence,
9 via G. Sansone 1, Sesto Fiorentino, 50019, Italy

10 ³Department of Physics, Università degli Studi di Genova and National Institute of Nuclear Physics INFN- Genoa, via
11 Dodecaneso 33, Genoa, 16146, Italy

12 ^anow at: Laboratory of Atmospheric Chemistry (LAC), Paul Scherrer Institut (PSI), Forschungsstrasse 111, Villigen,
13 5232, Switzerland

14

15 *Correspondence to: Roberta Vecchi (roberta.vecchi@unimi.it)

16

17 **Abstract.** In this paper, a new methodology coupling aerosol optical and chemical parameters in the same source
18 apportionment study is reported. In addition to results on sources contribution, this approach provides information such
19 as estimates for the atmospheric Absorption Ångström Exponent (α) of the sources and Mass Absorption Cross section
20 (MAC) for fossil fuel emissions at different wavelengths.

21 A multi-time resolution source apportionment study using Multilinear Engine ME-2 was performed on a PM10 dataset
22 with different time resolutions (24 hours, 12 hours, and 1 hour) collected during two different seasons in Milan (Italy) in
23 2016. Samples were optically analysed by a home-made polar photometer to retrieve the aerosol absorption coefficient
24 b_{ap} (in Mm^{-1}) at four wavelengths ($\lambda=405$ nm, 532 nm, 635 nm and 780 nm) and were chemically characterised for
25 elements, ions, levoglucosan, and carbonaceous components. The dataset joining chemically speciated and optical data
26 was the input for the multi-time resolution receptor model; this approach was proven to strengthen the identification of
27 sources thus being particularly useful when important chemical markers (e.g. levoglucosan, elemental carbon) are not
28 available. The final solution consisted in 8 factors (nitrate, sulphate, resuspended dust, biomass burning, construction

works, traffic, industry, aged sea salt); the implemented constraints led to a better physical description of factors and the bootstrap analysis supported the goodness of the solution. As for b_{ap} apportionment, consistently to what expected, biomass burning and traffic were the main contributors to aerosol absorption in the atmosphere. A relevant feature of the approach proposed in this work is the possibility of retrieving many other information about optical parameters; for example, opposite to the more traditional approach used by optical source apportionment models, here we obtained the source-dependent α value without any a priori assumption (α biomass burning = 1.83 and α fossil fuels = 0.80). In addition, MAC estimated for fossil fuel emissions was consistent with literature values.

It is worth noting that the here presented approach can be also applied using more common receptor models (e.g. EPA PMF instead of multi-time resolution ME-2) if the dataset comprises variables with the same time resolution as well as optical data retrieved by widespread instrumentation (e.g. an Aethalometer instead of home-made instrumentation).

39

40 **1. Introduction**

41 Atmospheric aerosol impacts both on local and global scale causing adverse health effects (Pope and Dockery, 2006), decreasing visibility (Watson, 2002), and influencing the climate (IPCC, 2013). To face these issues an accurate knowledge of aerosol emission sources is mandatory.

44 Currently, multivariate receptor models are considered a robust approach (Belis et al., 2015) to perform source apportionment studies and the Positive Matrix Factorization (PMF) (Paatero and Tapper, 1994) has become one of the most widely used receptor models (Hopke, 2016) in the aerosol community. In the late 1990s, the Multilinear Engine (ME-2) was developed and proven to be a very flexible algorithm to solve multilinear and quasi-multilinear problems (Paatero, 1999). The scripting feature of this algorithm allows the implementation of advanced receptor modelling approaches; one example is the multi-time resolution model (Zhou et al., 2004), which uses each experimental data in its original time schedule as model input. Source apportionment studies carried out by multi-time resolution model are still scarce in the literature (Zhou et al., 2004; Ogulei et al., 2005; Kuo et al., 2014; Liao et al., 2015; Crespi et al., 2016; Sofowote et al., 2018) although this methodology is very useful in measurement campaigns when instruments with different time resolutions (minutes, hours or days) are available as high time resolution data can be exploited without averaging them over the longest sampling interval.

55 It is noteworthy that the combination of time-resolved chemically speciated data with the information obtained from instrumentation measuring aerosol optical properties at different wavelengths (e.g. the absorption coefficient b_{ap}) is suggested as one of the future investigations of receptor modelling (Hopke, 2016); however, to the best of our knowledge, very few attempts in this direction have been done (e.g. Peré-Trepat et al., 2007; Xie et al., 2019). Wang et al. (2011, 2012) in a source apportionment study used the Delta-C ($\Delta C = BC@370 \text{ nm} - BC@880 \text{ nm}$ from Aethalometer

60 measurements) as an additional input variable and found that Delta-C was very useful in separating traffic from biomass
61 burning source contributions.

62 The wavelength dependence of b_{ap} can be empirically considered to be proportional to $\lambda^{-\alpha}$, where α is the Absorption
63 Ångström Exponent; α depends on particles composition and size, and it is a useful parameter to gain information about
64 particles type in atmosphere (see e.g. Yang et al., 2009). Among aerosol components, black carbon (BC) is the main
65 responsible for light absorption in atmosphere; in fact, it is considered the main aerosol contributor to global warming
66 and the second most important anthropogenic contributor after CO₂ (Bond et al., 2013). Black carbon refers to a fraction
67 of the carbonaceous aerosol characterised by peculiar features as for microstructure, morphology, thermal stability,
68 solubility, and light absorption (Petzold et al., 2013); in particular, it is characterised by a wavelength-independent
69 imaginary part of the refractive index over visible and near-visible regions. Another aerosol absorbing component is
70 brown carbon (BrC), referred to as light-absorbing organic matter with increasing absorption towards lower wavelengths,
71 especially in the UV region (Andreae and Gelencsér, 2006). BrC is an aerosol component that also affects the elemental
72 vs. organic carbon correct separation when using thermal-optical methods as outlined by Massabò et al. (2016).

73 Source apportionment models based only on multi-wavelength b_{ap} data are available in the literature, i.e. the widespread
74 Aethalometer model (Sandradewi et al., 2008a) and the more recent Multi-Wavelength Absorption Analyzer (MWA) model
75 (Massabò et al., 2015; Bernardoni et al., 2017b). Briefly, these models estimate the source contribution to aerosol
76 absorption exploiting their different dependence on λ (i.e. different α). As a step forward, MWA provides the b_{ap}
77 apportionment in relation to both the sources (i.e. fossil fuel combustion and biomass burning) and the components (i.e.
78 BC and BrC) and also provides an estimate for α of BrC. Indeed, source apportionment models based on optical data
79 usually assume two contributors to b_{ap} , namely fossil fuel combustion and biomass burning (only few exceptions are
80 present in the literature, e.g. Fialho et al., 2005). In most cases this assumption is well founded, except when episodic
81 events giving a not negligible contribution to aerosol absorption in the atmosphere occur, such as in presence of mineral
82 dust from the Saharan desert (Fuzzi et al., 2015). Moreover, the above-mentioned models need a priori assumptions on
83 the α values of the sources and wide ranges for α are reported in the literature (e.g. Sandradewi et al., 2008a); this is the
84 most critical step, since α depends on the kind of fuel, burning conditions and aging processes in the atmosphere. Without
85 accurate determination of source-specific atmospheric α (for example exploiting the information derived from source
86 apportionment using ¹⁴C measurements), the applicability of models based on optical data is questionable (Bernardoni et
87 al., 2017b; Massabò et al., 2015; Zotter et al., 2017). Moreover, the generally accepted assumption of $\alpha=1$ for fossil fuels
88 and BC, arising from the theory of absorption by spherical particles in the Rayleigh regime (Seinfeld and Pandis, 2006),
89 might not always be valid for aged atmospheric aerosol (Liu et al., 2018).

90 In the framework of a source apportionment study based on multi-time resolution receptor modelling, optical and chemical
91 datasets were joined to retrieve a multi- λ apportionment of b_{ap} with no need of a-priori assumptions on the contributing
92 sources. Instead of using α as an a priori input, this approach directly provided source-dependent α values. Moreover, the
93 multi- λ apportionment of b_{ap} in each source allowed to estimate MAC values at different wavelengths, exploiting the
94 well-known relation $EBC=b_{ap}(\lambda)/MAC(\lambda)$ (Bond and Bergstrom, 2006) where elemental carbon (EC) apportioned by the
95 model was considered as a proxy for BC. The evaluation of atmospheric MAC values is also not trivial due to the possible
96 presence of absorbing components different from BC (e.g. contribution from BrC, especially at lower wavelengths).
97 The original approach proposed in this work shows that coupling the chemical and optical information in a receptor
98 modelling process is particularly advantageous because: (1) strengthens the source identification that is particularly useful
99 when relevant chemical tracers (e.g. levoglucosan, EC) are not available; (2) gives estimates for source-specific
100 atmospheric α which are typically assumed a-priori in source apportionment models based on optical data; (3) provides
101 MAC values at different wavelengths for specific sources.

102 In this work, optical data were measured by a home-made multi-wavelength polar photometer and input data (chemical
103 + optical) in the receptor model comprised variables acquired with different time resolutions. Anyway, it is worth noting
104 that the here presented approach is of general interest as the same methodology could be applied to (1) datasets combining
105 aerosol chemical and optical data obtained by widespread instrumentation (e.g. Aethalometers for optical data); (2)
106 variables with the same time resolution.

107

108 **2. Material and methods**

109 *2.1 Site description and aerosol sampling*

110 Two measurement campaigns were performed during summertime (June-July) and wintertime (November-December)
111 2016 in Milan (Italy). Milan is the largest city (more than 1 million inhabitants, doubled by commuters everyday) of the
112 Po Valley, a very well-known hot-spot pollution area in Europe due to both large emissions from a variety of sources (i.e.
113 traffic, industry, domestic heating, energy production plants, and agriculture) and low atmospheric dispersion conditions
114 (e.g. Vecchi et al., 2007 and 2019; Perrone et al., 2012; Bigi and Ghermandi, 2014; Perrino et al., 2014).

115 The sampling site is representative of the urban background and it is situated at about 10 meters above the ground, on the
116 roof of the Physics Department of the University of Milan, less than 4 km far from the city centre (Vecchi et al., 2009).

117 It is important to note that during the sampling campaigns, a large building site was active next to the monitoring station.
118 Aerosol sampling was carried out using instrumentation with different time resolutions. Low time resolution PM10 data,
119 with a sampling duration of 24 and 12 hours during summertime (20 June-22 July 2016) and wintertime (21 November-

120 22 December 2016), respectively, were collected in parallel on PTFE (Whatman, 47 mm diameter) and pre-fired (700 °C,
121 1 hour) quartz-fibre (Pall, 2500QAO-UP, 47 mm diameter) filters. Low volume samplers with EPA PM10 inlet operating
122 at 1 m³ h⁻¹ were used. High time resolution data were collected during shorter periods (11 July-18 July and 21 November-
123 28 November 2016) by a streaker sampler (D'Alessandro et al., 2003). Shortly, the streaker sampler collects the fine and
124 coarse PM fractions (particles with aerodynamic diameter $d_{ac} < 2.5 \mu\text{m}$, and $2.5 < d_{ac} < 10 \mu\text{m}$, respectively) with hourly
125 resolution. Particles with $d_{ac} > 10 \mu\text{m}$ impact on the first stage and are discarded; the coarse fraction deposits on the second
126 stage, consisting of a Kapton foil; finally, the fine fraction is collected on a polycarbonate filter. The two collecting
127 supports are kept in rotation with an angular speed of about 1.8° h⁻¹ to produce a circular continuous deposit on both
128 stages.

129 Meteorological data were available at a monitoring station belonging to the regional environmental agency (ARPA
130 Lombardia) which is less than 1 km far away.

131

132 *2.2 PM mass concentration and chemical characterisation*

133 In this Section, chemical analyses performed on samples are summarised. As concentration detected in each sample was
134 characterised by its own uncertainty, only ranges for experimental uncertainties and minimum detection limits (MDLs)
135 for every set of variables are reported.

136 PM10 mass concentration was determined on PTFE filters by gravimetric technique. Weighing was performed by an
137 analytical balance (Mettler, model UMT5, 1 µg sensitivity) after a 24 hours conditioning period in an air-controlled room
138 as for temperature ($20 \pm 1 \text{ }^\circ\text{C}$) and relative humidity ($50 \pm 3 \%$) (Vecchi et al., 2004).

139 These filters were then analysed by Energy Dispersive X-Ray Fluorescence (ED-XRF) analysis to obtain the elemental
140 composition (details on the procedure can be found in Vecchi et al., 2004). For most elements and samples, concentrations
141 were characterised by relative uncertainties in the range 7-20 % (higher uncertainties for elements with concentrations
142 next to MDLs) and minimum detection limits of 0.9-30 ng m⁻³ with the above mentioned sampling conditions.

143 For each quartz-fibre filter, one punch (1.5 cm²) was extracted by sonication (1 h) using 5 ml ultrapure Milli-Q water and
144 levoglucosan and inorganic anions concentrations were quantified. Levoglucosan concentration was determined by High-
145 Performance Anion Exchange Chromatography coupled with Pulsed Amperometric Detection (HPAEC-PAD)
146 (Piazzalunga et al., 2010) only in winter samples. Indeed, as already pointed out by other studies at the same sampling
147 site (Bernardoni et al., 2011) and routinely assessed at monitoring stations in Milan by the Regional Environmental
148 Agency (private communication), levoglucosan concentrations during summertime are lower than the MDLs (i.e. about
149 6 ng m⁻³), due to both lower emissions (no influence of residential heating and negligible impact from other sources) and
150 higher OH levels in the atmosphere depleting molecular markers concentrations (Robinson et al., 2006; Hennigan et al.,

151 2010). Uncertainties on levoglucosan concentration were about 11 %. The quantification of main water-soluble inorganic
152 anions (SO_4^{2-} and NO_3^-) was performed by Ion Chromatography (IC); MDLs were 25 and 50 ng m^{-3} with summertime and
153 wintertime sampling conditions, respectively, and uncertainties were about 10 %. Unfortunately, due to technical
154 problems no data on ammonium were available. Details on the analytical procedure for IC analysis are reported in
155 Piazzalunga et al. (2013).

156 Another punch (1.0 cm^2) of each quartz-fibre filter was analysed by Thermal Optical Transmittance analysis (TOT, Sunset
157 Inc., NIOSH-870 protocol) (Piazzalunga et al., 2011) in order to assess organic and elemental carbon (OC and EC)
158 concentrations. MDLs were 75 and 150 ng m^{-3} with summertime and wintertime sampling conditions, respectively, and
159 uncertainties were in the range 10-15 %.

160 Hourly elemental composition was assessed by Particle Induced X-ray Emission (PIXE) technique, using a properly
161 collimated proton beam and scanning the deposits in steps corresponding to 1-hour aerosol deposit (details in Calzolari et
162 al., 2015). As low time resolution PM10 samples were also available, fine and coarse elemental concentrations determined
163 by PIXE analysis were added up to obtain PM10 concentrations with hourly resolution. PM10 hourly concentrations for
164 most elements and samples were characterised by relative uncertainties in the range 10-30 % (higher uncertainties for
165 elements near MDLs) and MDLs ranged from a minimum of 0.1 to a maximum of 15 ng m^{-3} (higher MDLs typically
166 detected for $Z < 20$ elements).

167

168 *2.3 Aerosol light-absorption coefficient measurements*

169 The aerosol absorption coefficient (b_{ap}) at the 4 wavelengths $\lambda = 405, 532, 635$ and 780 nm was measured on both low
170 and high time resolution samples with the home-made polar photometer PP_UniMI (Vecchi et al., 2014; Bernardoni et
171 al., 2017c).

172 Low time resolution optical measurements taken into account were those performed on PTFE filters since their physical
173 characteristics can be considered more similar to polycarbonate filters used by the streaker sampler. Moreover, previous
174 works reported a bias on b_{ap} measured by instrumentation using fibre filters (e.g. Cappa et al., 2008; Lack et al., 2008;
175 Davies et al., 2019; and references therein). Vecchi et al. (2014) found that b_{ap} at 635 nm was 40% higher when measured
176 on quartz-fibre filter compared to parallel samples collected on PTFE. This effect was ascribed to sampling artefacts due
177 to organics in aerosol samples collected in Milan.

178 As for high time resolution samples, b_{ap} was measured only in the fine fraction collected on polycarbonate filters since
179 absorption of the Kapton foil on which the coarse fraction was collected did not allow b_{ap} assessment. Anyway, b_{ap} values
180 in PM2.5 and PM10 were expected to be fairly comparable, as aerosol absorption in atmosphere is mostly due to particles
181 in the fine fraction at heavily polluted urban sites like Milan. To verify this assumption, high time resolution b_{ap} data in

182 PM2.5 were averaged over the time scale of low time resolution b_{ap} in PM10 and compared; the agreement was good,
183 between 11 % and 13 % depending on the λ , except for b_{ap} at $\lambda=405$ nm that showed a higher difference (27 %) but with
184 most data (83 %) within experimental uncertainties. To take into account this difference, b_{ap} data at $\lambda=405$ nm were
185 homogenised before using them in the model, following the criterion used for chemical species (for further detail about
186 homogenisation procedure, see Sect. 2.4 and Sect. 2.5).
187 Uncertainties on b_{ap} were quantified in 15 % and MDL was in the range 1-10 Mm^{-1} depending on sampling duration and
188 wavelength as already reported in Vecchi et al. (2014) and Bernardoni et al. (2017c). Pre-treatment procedure for
189 experimental uncertainties and MDLs was the same used for chemical variables in order to create suitable input matrices
190 required by the multi-time resolution model (see also Sect. 2.5). Optical system stability was checked during the
191 measurement session, evaluating the reproducibility of the measurement on a blank test filter. Laser stability was also
192 checked at least twice a day and the recorded intensities were used to normalise blank and sampled filters analysis.

193

194 *2.4 Model description*

195 Multivariate receptor models (Henry, 1997) are among the most widespread and robust approaches used to carry out
196 source apportionment studies for atmospheric aerosol (Belis et al., 2014 and 2015). In particular, the Positive Matrix
197 Factorization PMF2 (Paatero and Tapper, 1994; Paatero, 1997) had been extensively used in the literature and, afterwards,
198 the Multilinear Engine ME-2 (Paatero, 1999 and 2000) introduced the possibility of solving all kinds of multilinear and
199 quasi-multilinear problems. The fundamental principle of these modelling approaches is the mass conservation between
200 the emission source and the receptor site; using the information carried by aerosol chemical composition assessed in
201 samples collected at the receptor site, a mass balance analysis can be performed to identify the factors influencing aerosol
202 mass concentrations (Hopke, 2016). Factors can be subsequently interpreted as the main sources impacting the site,
203 through the knowledge about major sources in the investigated area and the exploitation of chemical fingerprints available
204 from previous literature works (Belis et al., 2014). Referring to the input data as matrix X (matrix elements x_{ij}), the
205 chemical profile of the factors as matrix F (matrix elements f_{kj}), and the time contribution of the factors as matrix G
206 (matrix elements g_{ik}), the main equation of a bilinear problem can be written as follows:

$$207 \quad x_{ij} = \sum_{k=1}^P g_{ik} f_{kj} + e_{ij} \quad (1)$$

208 where the indices i, j, and k indicate the sample, the species, and the factor, respectively; P is the number of factors and
209 the matrix E (matrix elements e_{ij}) is composed by the residuals, i.e. the difference between measured and modelled values.
210 In this way, a system of NxM equations is established, where N is the number of samples and M is the number of species.
211 The solution of the problem is computed minimising the object function Q defined as:

212
$$Q = \sum_{i=1}^N \sum_{j=1}^M \left(\frac{e_{ij}}{\sigma_{ij}} \right)^2 \quad (2)$$

213 where σ_{ij} are the uncertainties related to the input data.

214 The multi-time resolution receptor model was developed in order to use each data value in its original time schedule,
 215 without averaging the high time resolution data or interpolating the low time resolution data (Zhou et al., 2004; Ogulei et
 216 al., 2005). The main Eq. (1) is consequently modified as below:

217
$$x_{sj} = \frac{1}{t_{s2} - t_{s1} + 1} \sum_{k=1}^P f_{kj} \sum_{i=t_{s1}}^{t_{s2}} g_{ik} \eta_{jm} + e_{sj} \quad (3)$$

218 where the indices s, j, and k indicate the sample, the species and the factor respectively; P is the number of factors; t_{s1} and
 219 t_{s2} are the starting and ending time for the s-th sample in time units (i.e. the shortest sampling interval that is 1 hour for
 220 the dataset used here) and i represents one of the time units of the s-th sample. η_{jm} are adjustment factors for chemical
 221 species replicated with different time resolution and measured with different analytical methods (represented by the
 222 subscript m).

223 If η is close to unity, species concentration measured by different analytical approaches can be considered in good
 224 agreement; non-replicated species have adjustment factors set to unity by default. In this work, the adjustment factors
 225 were always set to unity in the model; to take into account the use two types of aerosol samplers (i.e. low volume sampler
 226 with EPA inlet and streaker sampler) and different analytical techniques to obtain the elemental composition (i.e. ED-
 227 XRF and PIXE), concentrations of replicated species with multiple time resolutions were homogenised before inserting
 228 them into the input matrix X, as will be explained in Sect. 2.5. This data treatment avoids the consistency check between
 229 η values calculated by the model and differences in experimental data characterised by high and low time resolution.
 230 Otherwise, this step should always be performed after running the model.

231 In the multi-time resolution model the following regularisation equation is introduced to take into account that some
 232 sources could contain few or no species measured with high time resolution:

233
$$g_{(i+1)k} - g_{ik} = 0 + \varepsilon_i \quad (4)$$

234 where ε_i represent the residuals.

235 As already pointed out by Ogulei et al. (2005), a weighing parameter for low resolution species might be necessary; in
 236 this study, it was implemented in the equations and set at 0.5 for strong species (not applied to weaker species such as
 237 Na, Mg, and Cr, see Sect. 2.5) in 24-h or 12-h samples.

238 Equations (3) and (4) are solved using the Multilinear Engine (ME) program (Paatero, 1999). In Eq. (2), the object
 239 function Q takes into account residuals from the main Eq. (3) and from the auxiliary equations (regularisation Eq. (4),
 240 normalisation equation, pulling equations, and constraints).

241 In this work, the multi-time resolution model implemented by Crespi et al. (2016) was used; therefore, constraints were
242 inserted in the model and the bootstrap analysis was also performed to evaluate the robustness of the final solution.

243

244 *2.5 Input data*

245 As already mentioned in Sect. 2.4, instead of using adjustment factors in the model (all set equal to one), concentrations
246 of replicated species with different time resolutions were pre-homogenised and then inserted into the input matrix X.
247 Concentration data with longer sampling interval (24 and 12 hours in this work) were considered as benchmark, since
248 analytical techniques usually show a better accuracy on concentration values far from MDLs (i.e. samples collected on
249 longer time intervals) (Zhou et al., 2004; Ogulei et al., 2005).

250 Variables were then classified as weak and strong according to the signal-to-noise ratio (S/N) criterion (Paatero, 2015).
251 For hourly data only strong variables ($S/N \geq 1.2$) were considered; for low time resolution data also weak variables such
252 as Na, Mg and Cr (with S/N equal to about 0.8) that resulted strong variables in hourly samples, were also included
253 although with associated uncertainties comparable to concentration values in order to avoid the exclusion of too many
254 data. Indeed, excluding these low time resolution variables from the analysis gave rise to artificial high values in the time
255 contribution matrix for sources traced by these species (in this case it was an issue for aged sea salt traced by Na and Mg,
256 see Sect. 3.2); this oddity was already reported by Zhou et al. (2004).

257 Every measured variable in each sample is characterised by its own uncertainty; ranges of experimental uncertainties and
258 MDLs are reported in Sect. 2.2 and 2.3 for chemical and optical analyses, respectively. Variables with more than 20 %
259 of the concentration data below MDL values were omitted from the analysis (Ogulei et al., 2005). Uncertainties, missing
260 values and data below minimum detection limits were pre-treated according to Polissar et al. (1998). In general, missing
261 concentration values were estimated by linear interpolation of the measured data and their uncertainties were assumed as
262 three times this estimated value (Zhou et al., 2004; Ogulei et al., 2005). As for summertime levoglucosan data (always
263 below MDLs), the approach was to include them as below MDL data and not as missing data following Zhou et al. (2004)
264 who underlined that the multi-time resolution model is more sensitive to missing values than the original PMF model. In
265 order to avoid double counting, in this study S was chosen as input variable instead of SO_4^{2-} as it was determined on both
266 low time and high time resolution samples (by XRF and PIXE analysis, respectively, see Calzolari et al., 2008). However,
267 elemental SO_4^{2-} and S concentrations showed a high correlation (correlation coefficient $R=0.98$) and the Deming
268 regression gave a slope of 2.69 ± 0.13 (sulphate vs. sulphur) with an intercept of $-198 \pm 82 \text{ ng m}^{-3}$, i.e. compatible with
269 zero within 3 standard deviations. The slight difference (of the order of 10%) between the estimated slope and the SO_4^{2-} -
270 to-S stoichiometric coefficient (i.e. 3) can be ascribed to either a small fraction of insoluble sulphate or to the use of
271 different analytical techniques.

272 PM10 mass concentrations were included in the model with uncertainties set at four times their values (Kim et al., 2003).
273 In the end, 22 low time resolution variables (PM10 mass, Na, Mg, Al, Si, S, K, Ca, Cr, Mn, Fe, Cu, Zn, Pb, EC, OC,
274 levoglucosan, NO_3^- , $b_{\text{ap}} 405\text{nm}$, $b_{\text{ap}} 532\text{nm}$, $b_{\text{ap}} 635\text{nm}$, $b_{\text{ap}} 780\text{nm}$) and 17 hourly variables (Na, Mg, Al, Si, S, K, Ca, Cr,
275 Mn, Fe, Cu, Zn, Pb, $b_{\text{ap}} 405\text{nm}$, $b_{\text{ap}} 532\text{nm}$, $b_{\text{ap}} 635\text{nm}$, $b_{\text{ap}} 780\text{nm}$) were considered.

276 The input matrix X consisted in 386 samples and the total number of time units was 1117. The analysis was performed
277 in the robust mode; lower limit for G contribution was set to -0.2 (Brown et al., 2015) and the error model $\text{em}=-14$ was
278 used for the main equation with $C_1=$ input error, $C_2= 0.0$, and $C_3=0.1$ (Paatero, 2012) for both chemical and optical
279 absorption data.

280 Sensitivity tests on the uncertainty of absorption data were performed starting from a minimum uncertainty of 10 %.
281 Lower uncertainties were considered not physically meaningful from an experimental point of view. ME-2 analyses
282 performed with 10 % uncertainty on absorption data gave very similar results to the base case solution presented in the
283 Supplement (Figure S1 and Table S3), with no differences in mass apportionment and a maximum variation in the
284 concentrations of chemical and optical profiles (matrix F) of 7 % when considering significant variables in each profile
285 (i.e. EVF higher or near 0.30). In contrary, considering an uncertainty of 20 % on absorption data, the solution
286 significantly differed from the one reported in the Supplement and showed less physical meaning (e.g. a couple of factors
287 got mixed, an additional unique factor appeared giving a null mass contribution). Thus, the estimated relative uncertainty
288 of 15 % was here considered appropriate for optical variables.

289 It is also noteworthy that ME-2/PMF analysis is not a-priori harmed by the use of joint matrices containing different units
290 (see e.g. Paatero, 2018). Indeed, if different units are present in different columns of matrix X, the output data in the factor
291 matrix G are pure numbers and elements in a column of the factor matrix F carry the same dimension and unit as the
292 original data in matrix X. In addition, the average total contribution to the mass of a specific source due to species in a
293 certain factor in matrix F must be retrieved a-posteriori summing up only mass contributions by chemical components
294 (i.e. excluding optical components in matrix F).

295 To the authors' knowledge, this was the first time that the absorption coefficient at different wavelengths was introduced
296 in the multi-time resolution model jointly with chemical variables and used to more robustly identify the sources;
297 moreover, this approach led to the assessment of source-dependent α and MAC values in an original way.

298

299 **3. Results and discussion**

300 *3.1 Concentration values*

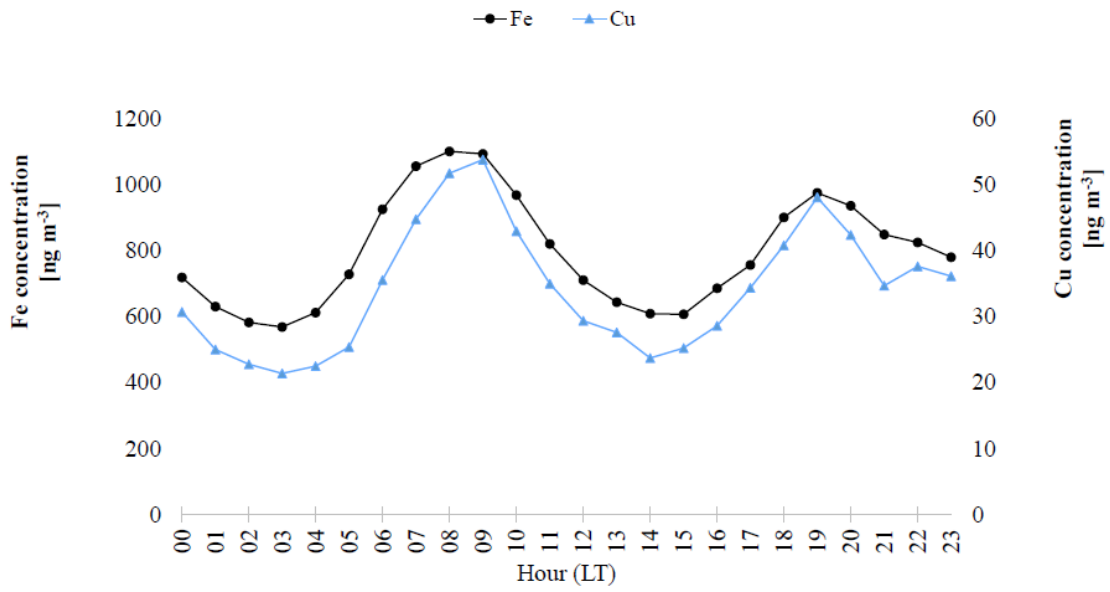
301 In Table S1 (Supplement) basic statistics on mass and chemical species concentrations at different time resolutions are
302 given.

303 Most variables showed higher mean and median concentrations during the winter campaign, when atmospheric stability
304 conditions influenced the monitoring site; exceptions were Al, Si and Ca which had lower median concentrations (as
305 detected in low time resolution samples). This was not unexpected as they are typical tracers of soil dust resuspension
306 (Viana et al., 2008) that can be more relevant during summertime due to drier soil conditions and stronger atmospheric
307 turbulence. Moreover, the good correlation between these elements (Al vs Si: $R^2=0.94$ and Ca vs. Si: $R^2=0.78$) suggested
308 their common origin.

309 Potassium showed the clearest seasonal behaviour in concentration values going from 284 ng m^{-3} ($10^{\text{th}}\text{-}90^{\text{th}}$ percentile:
310 $151\text{-}344 \text{ ng m}^{-3}$) to 660 ng m^{-3} ($10^{\text{th}}\text{-}90^{\text{th}}$ percentile: $349\text{-}982 \text{ ng m}^{-3}$) in summer and winter, respectively, in low time
311 resolution samples. K is an ambiguous tracer, since it is emitted by a variety of sources among which there are crustal
312 resuspension and biomass burning. In our dataset, wintertime K values showed a good correlation with levoglucosan
313 concentrations ($R^2=0.71$) suggesting the impact of biomass burning as levoglucosan is a well-known tracer for biomass
314 burning emissions in winter samples (Simoneit al., 1999). Also looking at K-to-Si ratio (where Si was taken as soil dust
315 marker) significant seasonal differences came out; it was 0.35 ± 0.15 in high time resolution summer samples and $2.0 \pm$
316 2.2 in winter ones, to be compared with the much more stable Al-to-Si ratio (i.e. 0.26 ± 0.04 and 0.28 ± 0.09 in summer
317 and winter, respectively) indicating a soil-related origin.

318 Among the elements typically associated to anthropogenic sources, Fe and Cu showed a good correlation (e.g. $R^2=0.72$
319 on hourly resolution samples) as well as Cu and EC (Cu vs EC: $R^2=0.84$, on low time resolution data); in addition, the
320 diurnal pattern of Fe and Cu showed traffic rush-hours peaks (7-9 a.m. and around 19 p.m. as shown in Fig.1). These
321 results were suggestive of a common source; in the literature these aerosol chemical components are reported as tracers
322 for vehicular emissions (e.g. Viana et al., 2008; Thorpe and Harrison, 2008).

323

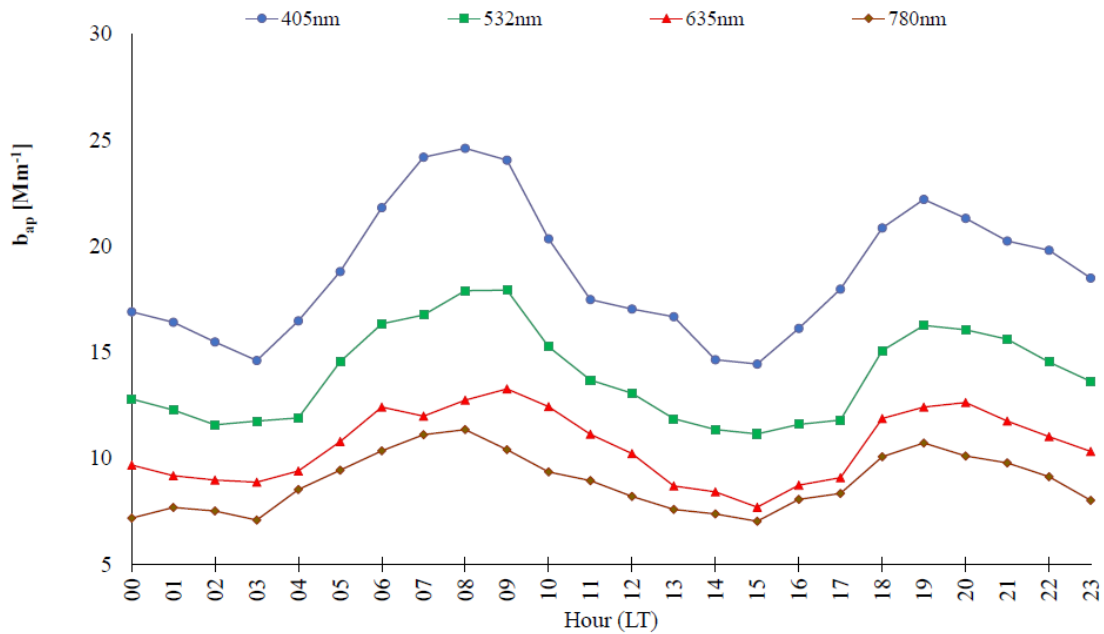


324

325 Figure 1: Diurnal profile of Fe and Cu concentrations (in ng m⁻³).

326

327 In Table S2 (Supplement) also basic statistics on b_{ap} values referred to low resolution samples collected on PTFE are
 328 reported. Diurnal mean temporal patterns for b_{ap} at different wavelengths (retrieved from hourly resolved data) are
 329 displayed in Fig. 2.



330

331 Figure 2: Diurnal profile of aerosol absorption coefficient (in Mm⁻¹) measured at different wavelengths.

332

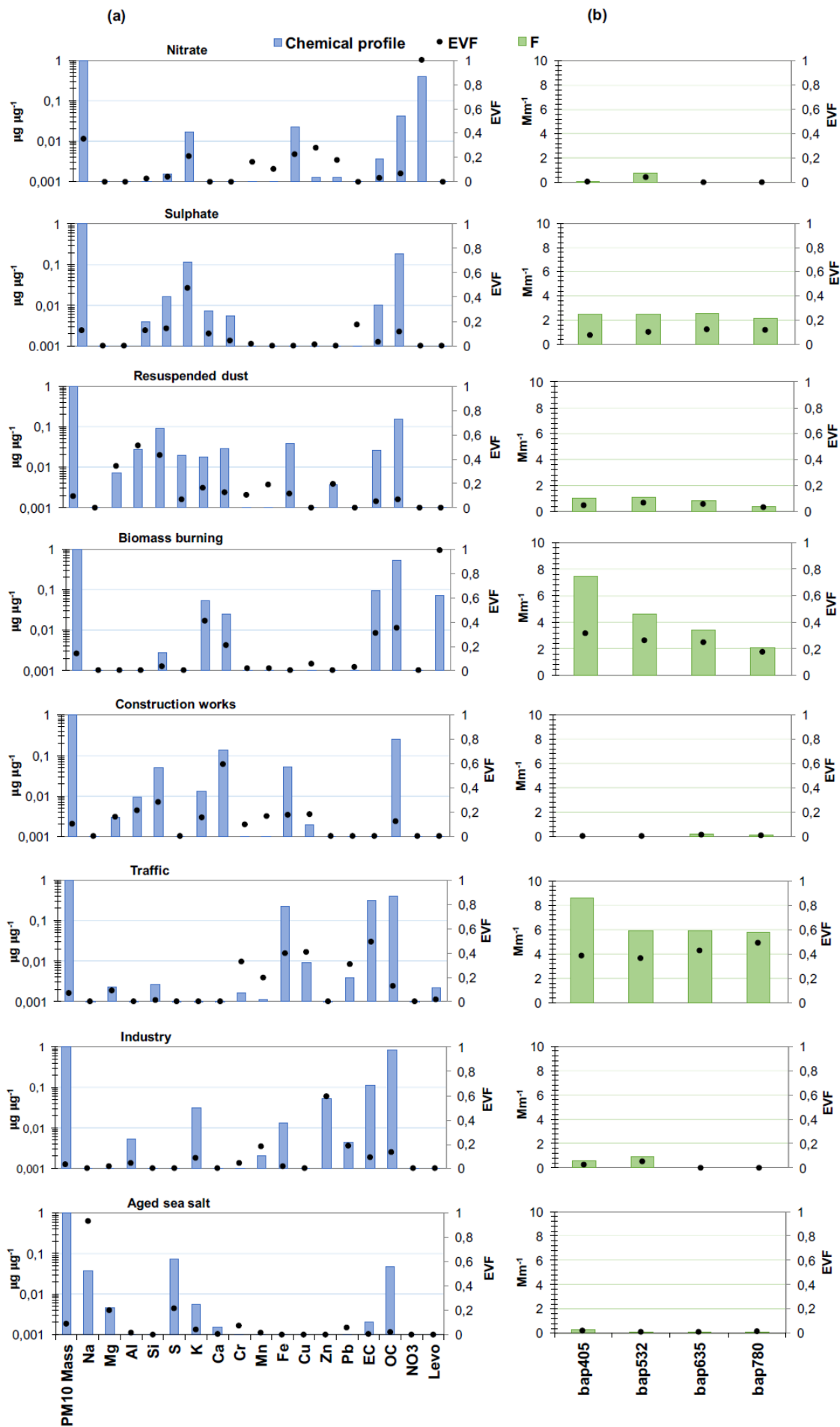
333 *3.2 Source apportionment with the multi-time resolution model*

334 Different solutions (from 5 to 10 factors) were explored; after 30 convergent runs, the 8-factor base-case solution
335 corresponding to the lowest Q value (2086.88) was firstly selected (see Fig. S1 in the Supplement). It is important to note
336 that the model was run using all variables (chemical + optical) as explained in Sect. 2.5. A lower or higher number of
337 factors caused ambiguous chemical profiles and the physical interpretation suggested clearly mixed sources for a lower
338 number of factors or unique factors in case of more factors (i.e. Pb for 9 factors); moreover, inconsistent mass closure
339 was detected increasing the number of factors (e.g. the sum of species contribution was up to 25 % higher than the mass
340 for the 10-factor solution). In the 8-factor base case solution, the mass was well reconstructed by the model ($R^2=0.98$),
341 with a slope of 0.98 ± 0.02 and negligible intercept ($0.51 \pm 0.89 \mu\text{g m}^{-3}$).

342 The factor-to-source assignment was based on both the Explained Variation for F matrix (EVF) values - which are
343 typically higher for chemical tracers (Lee et al., 1999; Paatero, 2010) - and the physical consistence of factor chemical
344 profiles. In the chosen solution, the not explained variation was lower than 0.25 for all variables. The uncertainty-scaled
345 residuals (Norris et al., 2014) showed a random distribution of negative and positive values in the ± 3 range, with a
346 Gaussian shape for most of the variables (Fig. S2 in the Supplement).

347 Using EVF and chemical profiles reported in Fig. S1(a), the 8 factors were tentatively assigned to nitrate, sulphate,
348 resuspended dust, biomass burning, construction works, traffic, industry, and aged sea salt. In Table S3 (in the
349 Supplement) absolute and relative average source contributions to PM10 mass are reported.

350 Although the above mentioned base-case solution was a satisfactory representation of the main sources active in the area
351 (as reported in previous works, see e.g. Marcazzan et al., 2003; Vecchi et al., 2009 and 2018; Bernardoni et al., 2011 and
352 2017a; Amato et al., 2016), the chemical profiles of some factors were improved exploring rotated solutions. The most
353 relevant case was represented by aged sea-salt where typical diagnostic ratios such as Mg/Na and Ca/Na (in bulk sea
354 water equal to 0.12 and 0.04, respectively, as reported e.g. in Seinfeld and Pandis, 2006) were not well reproduced in the
355 base-case solution and the chemical profile was too much impacted by the presence of Fe compared to bulk sea water
356 composition. Therefore, the above-mentioned diagnostic ratios were here used as constraints and Fe was maximally pulled
357 down in the chemical profile. The effective increase in Q was of about 61 units ($Q=2147$), with a percentage increase of
358 about 3 %; as a rule of thumb, an increase in the Q value of a few tens is generally considered acceptable (Paatero and
359 Hopke, 2009). It is noteworthy that the constrained solution led to an improvement in the chemical profile of the aged sea
360 salt and negligible differences in all other relevant features of the solution (i.e. EVF, residuals, mass reconstruction, source
361 apportionment) were found compared to the base-case solution. Therefore, the 8-factor constrained solution was
362 considered the most physically reliable; results are presented in Table 1 and Fig. 3 and discussed in detail in the following.



364 Figure 3: (a) Chemical profiles of the 8-factor constrained solution; (b) b_{ap} apportionment of the 8-factor constrained
 365 solution. The blue bars represent the chemical profile (output of the matrix F normalised on mass), the green bars the
 366 output values of the matrix F for the optical variables, and the black dots the EVF.

367

368

Factors	Summer [$\mu\text{g m}^{-3}$]	Winter [$\mu\text{g m}^{-3}$]	Total [$\mu\text{g m}^{-3}$]
Nitrate	3.6 (15 %)	21.1 (44 %)	10.2 (31 %)
Sulphate	6.3 (26 %)	8.1 (17 %)	7.0 (21 %)
Resuspended dust	4.6 (19 %)	1.7 (4 %)	3.5 (11 %)
Biomass burning	0.32 (1 %)	8.3 (17 %)	3.3 (10 %)
Construction works	5.9 (24 %)	3.4 (7 %)	4.9 (15 %)
Traffic	1.4 (6 %)	2.2 (5 %)	1.7 (5 %)
Industry	0.86 (4 %)	1.2 (3 %)	1.0 (3 %)
Aged sea salt	1.4 (6 %)	1.8 (4 %)	1.6 (5 %)

369

370 Table 1: Absolute and relative average source contributions to PM10 mass in the 8-factor constrained solution.

371

372 The factor interpreted as nitrate fully accounted for the explained variation of NO_3^- . This factor contained a significant
 373 fraction of nitrate in the chemical profile (39 %) and all nitrate was present only in this factor. This source was by large
 374 the most significant one at the investigated site, explaining about 31 % of the PM10 mass over the whole campaign (a
 375 similar estimate – 26 % - was reported by Amato et al. (2016) during the AIRUSE campaign in Milan in 2013) raising up
 376 to 44 % during wintertime (comparable to 37 % reported by Vecchi et al. (2018)). Indeed, the Po valley is well-known
 377 for experiencing very high nitrate concentrations during wintertime (Vecchi et al., 2018; and references therein) because
 378 of large emissions of gaseous precursors related to urban and industrial activities, biomass burning used for residential
 379 heating, high ammonia levels due to agricultural fields manure and poor atmospheric dispersion conditions.

380 The factor associated to sulphate showed $\text{EVF}=0.47$ for S and much lower EVF for all the other variables in the factor.
 381 Considering the sulphur contribution in the chemical profile in terms of sulphate and ammonium sulphate, the relative
 382 contribution of sulphur components in the profile increased from 11 % (S) up to 45 % (ammonium sulphate). The latter
 383 is the main sulphur compound detected in the Po valley as reported in previous papers such as e.g. Marcazzan et al. (2001)
 384 and was by far the highest contributor in the chemical profile. The other important contributor was OC (19 %), whose
 385 impact on PM mass increased up to 30 % when reported as organic matter using 1.6 as the organic carbon-to-organic
 386 matter conversion factor for this site (Vecchi et al., 2004). Due to the secondary origin of the aerosol associated to this
 387 factor, it was not surprising to find also a significant OC contribution; indeed, aerosol chemical composition in Milan is
 388 impacted by highly oxygenated components due to aging processes favoured by strong atmospheric stability (Vecchi et
 389 al., 2018 and 2019). In this factor, EC contributed for about 1 %. Considering the total EC concentration reconstructed

390 by the model, the EC fraction related to the sulphate factor was about 6 %. Opposite to sulphates, EC has a primary origin;
391 however, its presence with a very similar percentage (4-5 %) in a sulphate chemical profile was previously pointed out in
392 Milan, indicating a more complex mixing between primary and secondary sources (Amato et al., 2016) e.g. with sulphate
393 condensation on primary emitted particles. The sulphate factor accounted for 21 % of the PM10 mass.

394 The factor identified as resuspended dust is mainly characterised by high EVF and contributions coming from Al, Si and
395 Mg, i.e. crustal elements. The Al/Si ratio is 0.31, very similar to the literature value for average crustal composition
396 (Mason, 1966); the relatively high OC contribution in the chemical profile (15 %) together with the presence of EC (about
397 2.6 %) is suggestive of a mixing with road dust (Thorpe and Harrison, 2008). This source accounts for about 11 % of the
398 PM10 mass.

399 The factor identified as biomass burning was characterised by high EVF for levoglucosan (0.98), a known tracer for this
400 source as it is generated by cellulose pyrolysis; EVF higher than 0.3 were also found for K, OC, and EC. In the source
401 chemical profile, OC contributed for 54 %, EC for 10 %, levoglucosan for 7 %, and K for 5 %. The average biomass
402 burning contribution during this campaign was 10 % (up to 17 % in wintertime). Anticipating the discussion presented in
403 detail in Sect. 3.3, it is worth noting that the second largest contribution to the aerosol absorption coefficient after traffic
404 was detected in this factor.

405 The factor with high EVF (0.60) for Ca was associated to construction works, following literature works (e.g. Vecchi et
406 al., 2009; Bernardoni et al., 2011; Dall'Osto, 2013; Crilley et al., 2017; Bernardoni et al., 2017a; and references therein).
407 Major contributors to the chemical profile were Ca (13 %), OC (26 %), Fe, and Si (5 % each). This factor accounted on
408 average for 15 % to PM10 mass. As already mentioned, during the campaign a not negligible contribution from this
409 source was expected, due to the presence of a building site nearby the monitoring location.

410 In the factor assigned to traffic (primary contribution), EVF larger than 0.3 characterised EC, Cu, Fe, Cr, and Pb. The
411 highest relative mass contributions in the chemical profile were given by OC (41 %), EC (32 %), Fe (23 %), and Cu (1
412 %). The lack of relevant crustal elements such as Ca and Al in the chemical profile, suggested a negligible impact of road
413 dust in this factor. As reported above, at our sampling site the road dust contribution was very likely mixed to resuspended
414 dust and further separation of these contributions was not possible. This traffic (primary) contribution over the whole
415 dataset accounted for 5 % of the PM10 mass with a slightly lower absolute contribution in summer (see Table 1). This
416 contribution is comparable to the percentage (7 %) reported by Amato et al. (2016) for exhaust traffic emissions but it is
417 lower than our previous estimates (Bernardoni et al., 2011; Vecchi et al., 2018), i.e. 15 % in 2006 in PM10 and 12 % in
418 PM1 recorded in winter 2012. However, the current estimate seems to be still reasonable when considering the efforts
419 done in latest years to reduce vehicles exhaust particle emissions and the fraction of secondary nitrate due to high nitrogen
420 oxides and ammonia emissions in the region (INEMAR ARPA-Lombardia, 2018) which has to be added to account for

421 the overall traffic impact). Unfortunately, the non-linearity of the emission-to-ambient concentration levels relationship
422 and the high uncertainties in emission inventories still prevent a robust estimate of this secondary contribution to total
423 traffic exhaust emissions. As shown in Sect. 3.3, traffic is the largest contributor to aerosol absorption coefficient thus
424 strengthening the interpretation of this factor as a traffic emission source.

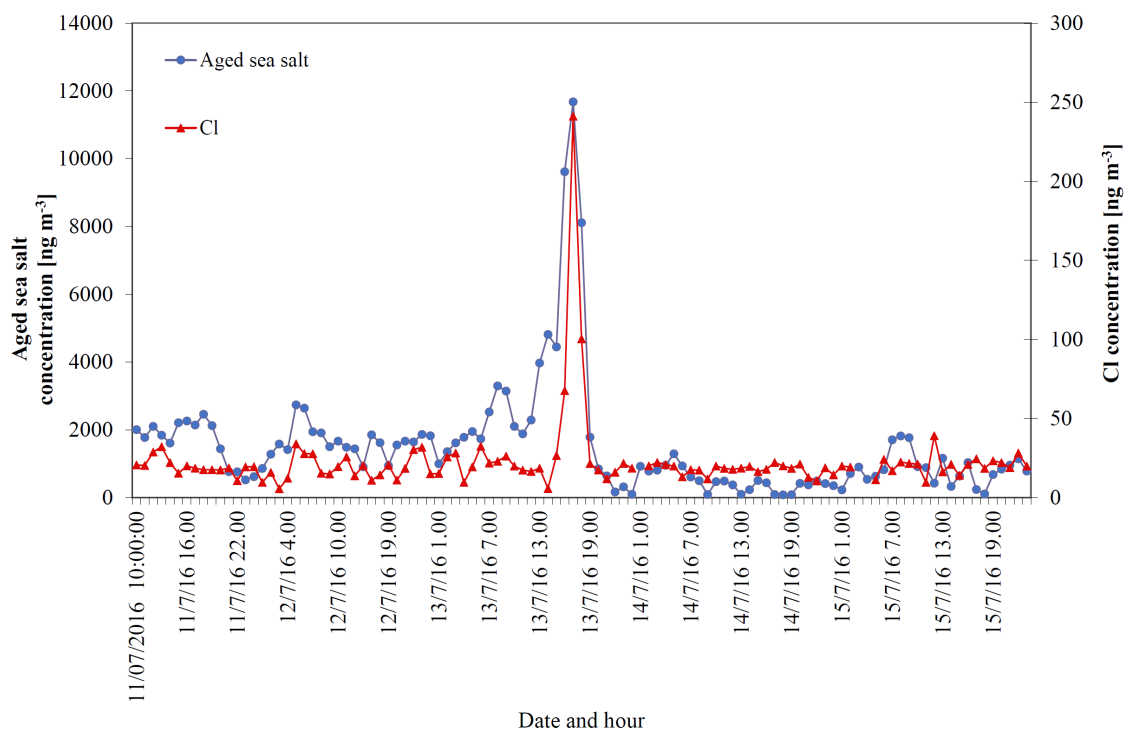
425 The industry factor showed high EVF for Zn (0.59) and the second highest EVF was related to Mn (0.13). Previous studies
426 at the same sampling site identified these elements as tracers for industrial emissions (e.g. Vecchi et al., 2018; and
427 references therein). The chemical profile resulted enriched by heavy metals and, after traffic, it was the profile with the
428 highest share of Cr, Mn, Fe, Cu, Zn, and Pb (explaining about 8 % of the total PM₁₀ mass in the profile). The industry
429 contribution was not very high in the urban area of Milan, accounting for 3 % on average.

430 The factor interpreted as aged sea salt was characterised by high EVF of Na (0.93) and this element was - as a matter of
431 fact - present only in this factor chemical profile. To check the physical consistency of this assignment and considering
432 that Milan is about 120 km away from the nearest sea coast, back-trajectories coloured by the aged sea salt concentration
433 (in ng m⁻³) were calculated through the NOAA HYSPLIT trajectory model (Draxler and Hess, 1998; Stein et al., 2015;
434 Rolph et al., 2017) and represented using the R package (Carslaw and Ropkins, 2012; R Core Team, 2019). As an example,
435 results from a very short event (13/07 h.16-18) singled out by the model and representing the highest sea salt contribution
436 during summer are reported in Fig. S3 (Supplement). Before and during the event, south-western air masses coming from
437 the Ligurian sea were observed while soon after the event, there was a rapid change of wind direction. High wind speeds
438 were recorded during the episode ($4.8 \pm 1.7 \text{ m s}^{-1}$ with a maximum peak of 9.5 m s^{-1}) compared to $1.9 \pm 1.0 \text{ m s}^{-1}$ average
439 wind speed characterising the summer campaign.

440 When marine air masses are transported to polluted sites, sea salt particles show a Cl deficit due to reactions with sulphuric
441 and nitric acid (Seinfeld and Pandis, 2006) and the factor chemical profile is expected to be enriched in sulphate and
442 nitrate. In this work, nitrate was not present in the aged sea salt chemical profile; a very rough estimate (Lee et al., 1999)
443 gave a maximum expected contribution of 2 % (about 82 ng m^{-3}) of the total nitrate mass in atmosphere that can be
444 considered negligible in terms of mass contribution of the sources.

445 Temporal patterns of Cl concentrations (not inserted in the multi-time resolution analysis as being a weak variable) during
446 marine aerosol episodes were exploited to further confirm the factor-to-source association. Cl concentration and aged sea
447 salt pattern showed an evident temporal coincidence in peak occurrence during the short summer event (Fig. 4), thus
448 supporting the source identification. Moreover, during this episode only the Cl coarse fraction increased (Fig. S4, in the
449 Supplement) and reached about 90 % of total PM₁₀ Cl concentration; Cl/Na ratio was 0.38 ± 0.05 , consistent with an
450 aging of marine air masses during advection showing the typical Cl depletion.

451



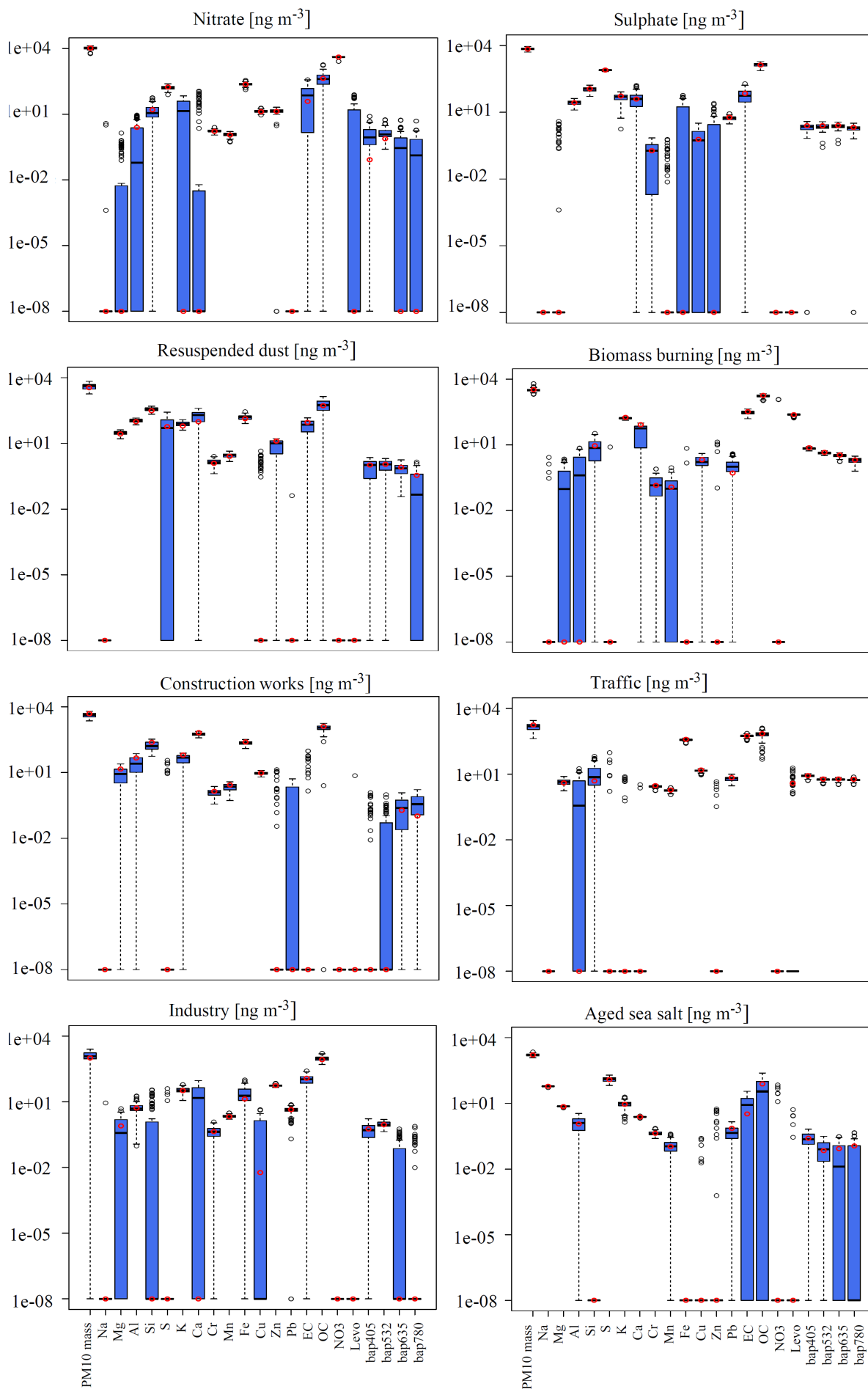
452

453 Figure 4: Temporal patterns of aged sea salt source retrieved from the multi-time resolution model and Cl concentrations
 454 measured in atmospheric aerosol.

455

456 Bootstrap analysis was performed to evaluate the uncertainties associated to source profiles (Crespi et al., 2016). 100 runs
 457 were carried out (see Fig. 5, values expressed in ng m⁻³ or Mm⁻¹ on a logarithmic scale); factors were well mapped, with
 458 Pearson coefficients always higher than 0.97, and tracers for each source showed small interquartile range, supporting the
 459 goodness of the solution presented in this work.

460



462 Figure 5: Box plot of the bootstrap analysis on the 8-factor constrained solution. The red dots represent the output values
463 of the solution of the model; the black lines the medians from the bootstrap analysis; the blue bars the 25th and 75th
464 percentile; the dotted lines the interval equal to 1.5 the interquartile range and the black dots the outliers from this interval.

465

466 *3.3 Improving source apportionment with optical tracers*

467 First of all, the use of the absorption coefficient determined at different wavelengths as input variable in the multi-time
468 resolution model, strengthened the identification of the sources suggesting that it can be exploited when specific chemical
469 tracers are not available (e.g. levoglucosan for biomass burning). To prove that, a separate source apportionment study
470 was performed with EPA PMF 5.0 (Norris et al., 2014) using only hourly elemental concentrations from samples collected
471 by the streaker sampler and hourly b_{ap} at different λ measured by PP_UniMI on the same filters as input variables. Streaker
472 samples typically lack of a complete chemical characterisation; in particular, important chemical tracers such as
473 levoglucosan and EC are not available. In this analysis, b_{ap} assessed at different wavelengths resulted effective in
474 identifying the biomass burning factor that explained a significant percentage of the b_{ap} itself (from 25 % to 35 %
475 depending on λ) (Fig. S5, in the Supplement); without the optical variables, the factor-to-source assignment would be
476 otherwise based only on the presence of elemental potassium although it is well-known that K cannot be considered an
477 unambiguous tracer as it is emitted by a variety of sources (see for example Pachon et al., 2013; and references therein).
478 Furthermore, results showed that the absorption coefficient contribution was higher than 45 % in the factor labelled as
479 traffic, highlighting the importance of exhaust emissions in a factor that would be differently characterised by elements
480 related to non-exhaust emissions (Cu, Fe, Cr).

481 From the multi-time resolution model, the two factors identified as biomass burning and traffic were the main contributors
482 to aerosol absorption in atmosphere and showed significant EVF values. At 780 and 405 nm, traffic contributions to b_{ap}
483 were 55 % and 42 %; at the same wavelengths biomass burning accounted for 20 % and 36 %. The Explained Variation
484 (EVF) of b_{ap} has the maximum value at 405 nm for biomass burning (0.32) and at 780 nm for traffic (0.49), showing the
485 tendency to decrease and increase with the wavelength, respectively.

486 The third contributor to aerosol absorption in atmosphere was the sulphate factor, with a contribution comparable to the
487 biomass burning one at 780 nm (about 20 % of the total reconstructed b_{ap} at this wavelength). The sulphate factor
488 contained a small fraction of EC, as previously discussed (see Sect. 3.2). This might be explained considering that
489 non/weakly light-absorbing material can form a coating able to enhance particle absorption (Bond and Bergstrom, 2006;
490 Fuller et al., 1999) within a few days after emission. Laboratory experiments and simulations from in-situ measurements
491 highlighted absorption amplification for absorbing particles coated with secondary organic aerosol (Schnaiter et al., 2003;
492 Moffet and Prather, 2009). Particle aging is a significant process in the Po valley due to low atmospheric dispersion

493 conditions and it might explain the relatively high contribution of the sulphate factor to the absorption coefficient in
494 respect to the other sources (apart from traffic and biomass burning). Resuspended dust was the main contributor at all
495 wavelengths (between 3 % and 7 % of the total reconstructed b_{ap} , depending on the wavelength), likely due to the role of
496 iron minerals. The other sources were less relevant in terms of EVF values and overall contributed for less than 11 %.

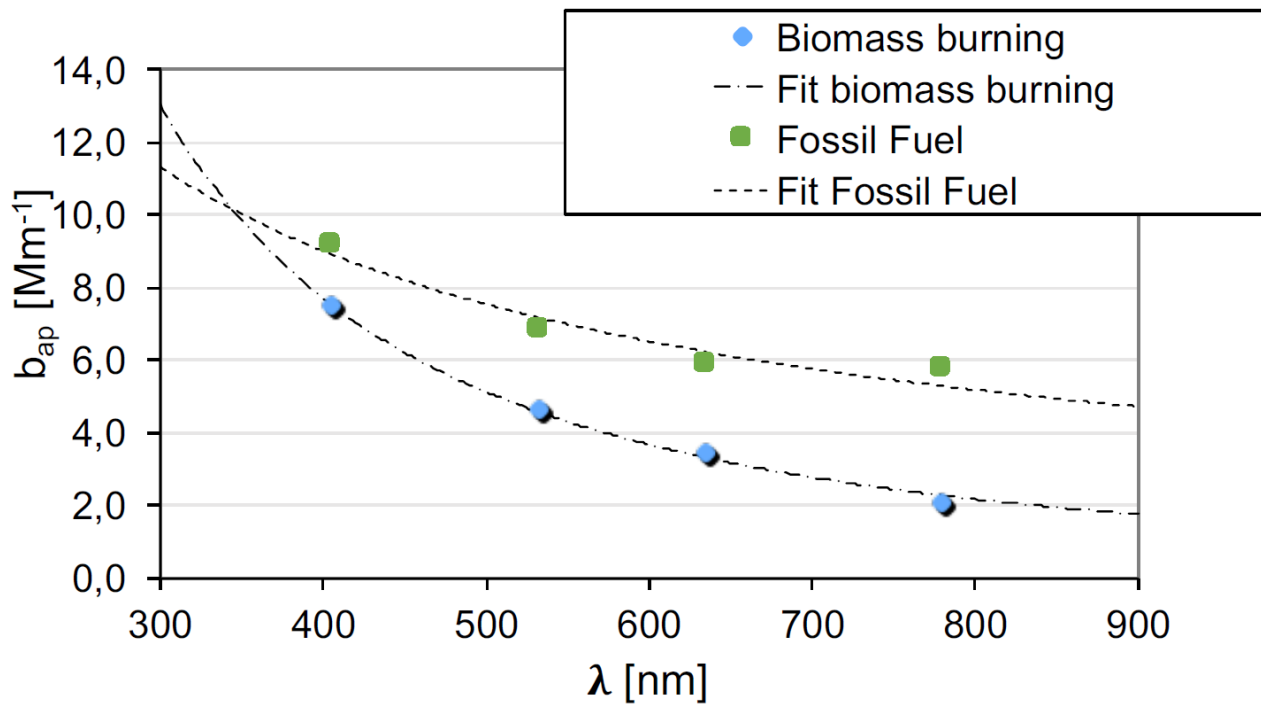
497 In contrast to the approach used in source apportionment models based on optical data like the widespread Aethalometer
498 model (Sandradewi et al., 2008a) and MWAA model (Massabò et al., 2015; Bernardoni et al., 2017b), it is noteworthy
499 that no a-priori information about α values of the fossil fuel and biomass burning sources was introduced in the multi-
500 time resolution model and an estimate for the values was directly retrieved from the model. Another literature approach
501 used Delta-C as an input variable together with chemical aerosol components in source apportionment models and was
502 very effective in separating traffic (especially diesel) emissions from biomass combustion emissions (Wang et al., 2011,
503 2012).

504 In order to compare the multi-time resolution model and models based on optical data, contributions due to traffic and
505 industry (i.e. emissions most likely connected to fossil fuel usage) were added up and labelled as “fossil fuel emissions”.

506 In accordance with the two-source approach used in the Aethalometer model, the discussion about optical properties will
507 be hereafter focused on the biomass burning and fossil fuel sources considering that sulphate and resuspended dust factors
508 were less significant also in terms of EVF for optical variables, ranging from 0.08 to 0.12 and from 0.03 and 0.06,
509 respectively, depending on the wavelength.

510 In Fig. 6 the wavelength dependence of b_{ap} for the biomass burning and the fossil fuel profiles obtained with the multi-
511 time resolution model is shown; as α values can show significant differences when calculated using different pairs of λ
512 (Sandradewi et al., 2008b), here we performed a fitting procedure considering $b_{ap} \propto \lambda^{-\alpha}$. Results were α_{BB} (α biomass
513 burning) = 1.83 and α_{FF} (α fossil fuels) = 0.80; the range of variability of α values was estimated with the bootstrap
514 analysis obtaining 0.78-0.88 for α_{FF} and 1.65-1.88 for α_{BB} (as 25th and 75th percentile, respectively).

515



516

517 Figure 6: b_{ap} dependence on λ for biomass burning and fossil fuel emissions.

518

519 Zotter et al. (2017) reported a possible combination of $\alpha_{FF}=0.8$ and $\alpha_{BB}=1.8$ when EC concentration from fossil fuel
 520 combustion (estimated with radiocarbon measurements) is between 40 % and 85 % of the total EC concentration; in this
 521 work, the fraction of EC ascribed by the multi-time model to fossil fuel sources was 56 %. The combination 0.9 and 1.68
 522 for α_{FF} and α_{BB} , respectively, was also suggested when in the study there are no or only limited additional information
 523 (e.g. from ¹⁴C measurements). From the wide range of possible combinations reported in the literature it is clear that the
 524 assessment of α_{BC} (assumed to be equal to α_{FF} in source apportionment models based on optical data) is still an issue and
 525 both experimental and simulation studies are in progress to reduce uncertainties and give a better evaluation of this key
 526 parameter.

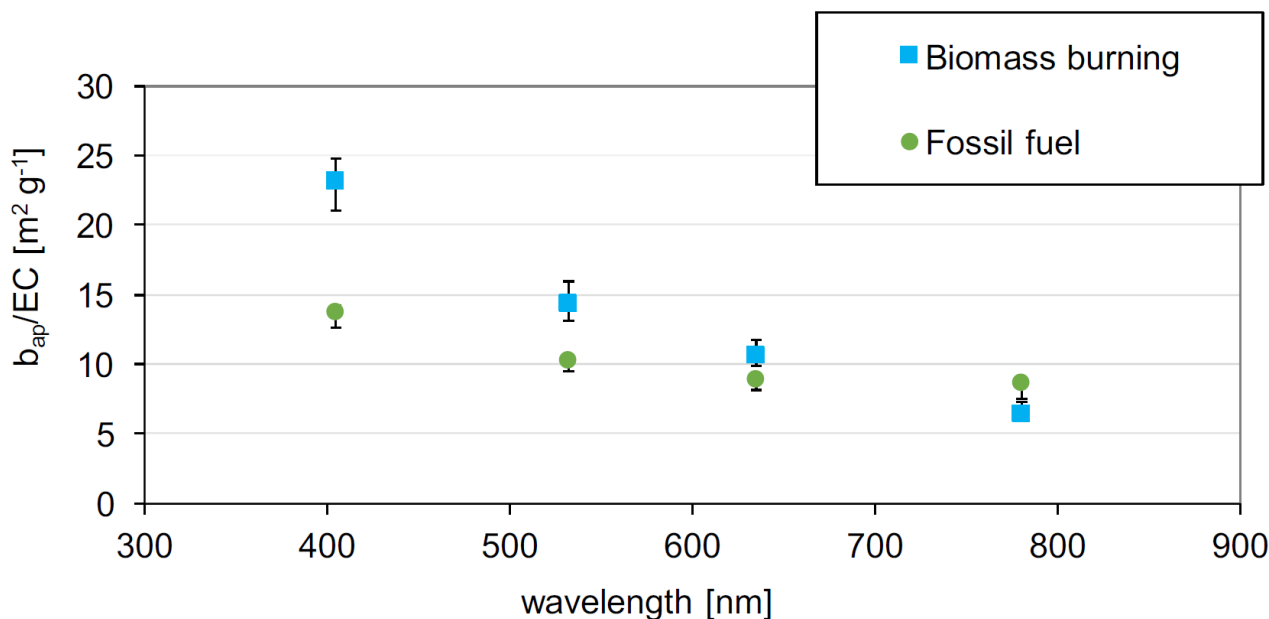
527 The α_{BB} value retrieved by the model was very similar to values reported by Zotter et al. (2017) and also comparable to
 528 1.86 found for biomass burning by Sandradewi et al. (2008a) and 1.8 obtained by Massabò et al. (2015) who used also
 529 independent ¹⁴C measurements for checking. The α_{FF} value resulted in the range 0.8-1.1 typically reported in source
 530 apportionment studies based on optical data (e.g. Bernardoni et al., 2017b; Zotter et al., 2017; and references therein).
 531 Indeed, the sampling site was an urban background station in Milan and our samples were hardly impacted by fresh traffic
 532 emissions. Considering the aged nature of Milan aerosol, the average α_{FF} was comparable to estimates for BC coated
 533 particles reported in the literature (approx. 0.6-1.3, see e.g. Liu et al., 2018) and obtained by both ambient measurement

534 (e.g. Fischer and Smith, 2018; and references therein) and numerical simulations (e.g. Gyawali et al., 2009; Liu et al.
535 2018; and references therein).

536 Results here reported allow also to study the relationship between the absorption coefficient and the mass of black carbon
537 (BC), i.e. the so called Mass Absorption Cross section at different wavelengths. The $MAC(\lambda) = b_{ap}(\lambda)/BC$ relationship
538 assumes that BC is the only light-absorbing species present; however, this assumption is not always valid since the
539 transport of mineral dust from desert areas and brown carbon can significantly contribute to aerosol absorption. During
540 our monitoring campaign, no contribution from Saharan dust was observed; opposite, biomass burning was proven to be
541 an important source so that BrC was certainly a significant contributor (Fuzzi et al., 2015) as also suggested by $\alpha_{BB} = 1.83$
542 in the biomass burning factor. The possible overestimation of BC when total b_{ap} is ascribed to BC only is usually
543 minimised choosing a wavelength higher than 600 nm, exploiting the spectral dependence of absorption from different
544 aerosol compounds (Petzold et al., 2013).

545 EC concentration retrieved from the chemical profiles (see Fig. 3) was used as a proxy for BC to estimate source-
546 dependent $b_{ap}(\lambda)$ -to-BC ratio. Results are represented in Fig. 7. It is noteworthy that here this ratio is intentionally not
547 indicated as MAC, since overestimation of the BC absorption especially at lower λ might occur (see previous discussion).
548 BrC is expected to give a small contribution in the fossil fuel source; therefore, the best approximation for $MAC(\lambda)$ values
549 are likely the $b_{ap}(\lambda)$ -to-BC ratios observed in the fossil fuel source at our monitoring site. They resulted $13.7 \text{ m}^2 \text{ g}^{-1}$ for λ
550 = 405 nm, $10.2 \text{ m}^2 \text{ g}^{-1}$ for $\lambda = 532 \text{ nm}$, $8.8 \text{ m}^2 \text{ g}^{-1}$ for $\lambda = 635 \text{ nm}$, $8.6 \text{ m}^2 \text{ g}^{-1}$ for $\lambda = 780 \text{ nm}$. For $\lambda = 550 \text{ nm}$ Bond and
551 Bergstrom (2006) reported a MAC value of $7.5 \pm 1.2 \text{ m}^2 \text{ g}^{-1}$ for uncoated fresh emitted particles and MAC values in
552 polluted regions ranging from 9 to $12 \text{ m}^2 \text{ g}^{-1}$, attributable to absorption enhancement due to particles coating. The MAC
553 estimate obtained in this work from multi-time resolution model for 532 nm is comparable to literature values above
554 reported thus confirming the importance of aging processes in atmosphere on the optical properties of particles.

555



556

557

558 Figure 7: b_{ap} -to-EC ratio dependence on λ for biomass burning and fossil fuel emissions. Error bars represent the 25th and
559 75th percentile retrieved from the bootstrap analysis.

560

561 Ratios represented in Fig. 7 are less comparable at $\lambda=405$ nm (see also Table S4, in the Supplement) due to the significant
562 contribution of BrC to b_{ap} at this wavelength in the biomass burning factor.

563 No seasonal differences in the atmospheric ratios were observed but at $\lambda = 405$ nm (see Table S4, in the Supplement), for
564 which winter values were higher than summer ones (17.8 ± 0.4 and 14.2 ± 0.5 , respectively); due to the influence of
565 biomass burning emissions on BrC concentration in atmosphere during the cold season.

566 From the outputs of the modelling approach here proposed, the apportionment of the biomass burning and fossil fuel
567 contributions to b_{ap} at different wavelengths was also obtained. As expected, the relative contribution to the total
568 reconstructed b_{ap} ascribed to the biomass burning factor decreased with increasing λ , opposite to the contribution from
569 fossil fuel combustion which gave the highest contribution at 780 nm (Table 2); in addition, the latter contribution
570 prevailed at all wavelengths at the investigated site.

571

	$\lambda = 405$ nm	$\lambda = 532$ nm	$\lambda = 635$ nm	$\lambda = 780$ nm
Biomass burning	36 % (31 %-36 %)	29 % (25 %-30 %)	26 % (23 %-27 %)	20 % (16 %-22 %)
Fossil fuels	45 % (41 %-46 %)	43 % (39 %-44 %)	45 % (41 %-47 %)	55 % (48 %-55 %)

572

573 Table 2: Average contribution to total reconstructed b_{ap} for the biomass burning and fossil fuel factors; in parenthesis 25th
574 and 75th percentile are reported.

575

576 **4. Conclusions**

577 The multi-time resolution model implemented through the Multilinear Engine (ME-2) script allowed the analysis of
578 experimental data collected at different time scales, coupling the detailed chemical speciation at low time resolution and
579 the temporal information given by high time resolution samples. The effect of the introduction of the aerosol absorption
580 coefficient (b_{ap}) measured at different wavelengths in the modelling process was investigated and gave promising results.
581 First of all, a more robust identification of sources was provided; secondly, it paved the way to the retrieval of optical
582 apportionment and optical characterisation of the sources (e.g. estimate of source-specific Absorption Ångström Exponent
583 - α - and Mass Absorption Cross section – MAC - at different wavelengths). It is worthy to note that currently in source
584 apportionment models based on optical data (e.g. Aethalometer model) values for α related to fossil fuel emissions and
585 biomass burning are fixed by the modeller thus carrying a large part of the uncertainties in the model results. Considering
586 that the estimates for the Absorption Ångström Exponent were here obtained as a result of a quite complex modelling
587 approach (i.e. using multi-time resolution datasets joining chemical and optical variables) and without any a-priori
588 assumption, the results obtained were fairly comparable to literature results and gave a further tool to assess more robust
589 source-related α values. Obviously these estimates are affected by a certain degree of uncertainty due to both experimental
590 data and modelling process (while uncertainties are typically not taken into consideration for fixed α values used in the
591 literature). In perspective, joining together different approaches such as the receptor modelling here proposed and e.g. ¹⁴C
592 data and artefact-free b_{ap} measurements will lead to better estimates of the Absorption Ångström Exponent.

593 The original approach described in this work can be applied to source apportionment studies using any suitable dataset
594 (not necessarily with multi-time resolution). Besides the traditional source apportionment, the impact of different sources
595 on the aerosol absorption coefficient was estimated; this piece of information can be very useful to formulate strategies
596 of pollutants abatement, in order to improve air quality and to face climate challenges. In particular, at the investigated
597 site secondary compounds constituted the highest contribution in terms of PM10 mass (52 % on average), while the two
598 factors identified as biomass burning and traffic were found to be the most significant contributors to aerosol light
599 absorption in atmosphere, in agreement with available literature.

600

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608

609 **Data availability.**

610 The data in the study are available from the authors upon request (roberta.vecchi@unimi.it).

611

612 **Supplement.**

613 The supplement related to this article is available online.

614

615 **Author contributions.**

616 ACF performed streaker sampling and related optical analysis, implemented the advanced model, analysed the results,
617 and drafted the paper. GV contributed to model implementation, data reduction and Hysplit back-trajectories retrieval.
618 VB, SV, and REP carried out the sampling campaign on filters, performed the optical measurements and data analysis.
619 GC, SN, and FL performed PIXE analysis and data reduction. DM and PP carried out ionic characterisation on filters and
620 data analysis. RV was responsible for the design and coordination of the study, the synthesis of the results and the final
621 version of the paper. All authors contributed to the interpretation of the results obtained with the new approach here
622 described and revised the manuscript content giving a final approval of the version to be submitted. RV and ACF reviewed
623 the paper addressing reviewers' comments.

624

625 **Competing interests.**

626 The authors declare that they have no conflict of interest.

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882 **List of Figure and Table Captions**

883 Figure 1: Diurnal profile of Fe and Cu concentrations (in ng m^{-3}).

884 Figure 2: Diurnal profile of aerosol absorption coefficient (in Mm^{-1}) measured at different wavelengths.

885 Figure 3: (a) Chemical profiles of the 8-factor constrained solution; (b) b_{ap} apportionment of the 8-factor constrained
886 solution. The blue bars represent the chemical profile (output of the matrix F normalised on mass), the green bars the
887 output values of the matrix F for the optical variables, and the black dots the EVF.

888 Figure 4: Temporal patterns of aged sea salt source retrieved from the multi-time resolution model and Cl concentrations
889 measured in atmospheric aerosol.

890 Figure 5: Box plot of the bootstrap analysis on the 8-factor constrained solution. The red dots represent the output values
891 of the solution of the model; the black lines the medians from the bootstrap analysis; the blue bars the 25th and 75th
892 percentile; the dotted lines the interval equal to 1.5 the interquartile range and the black dots the outliers from this interval.

893 Figure 6: b_{ap} dependence on λ for biomass burning and fossil fuel emissions.

894 Figure 7: b_{ap} -to-EC ratio dependence on λ for biomass burning and fossil fuel emissions. Error bars represent the 25th and
895 75th percentile retrieved from the bootstrap analysis.

896

897 Table 1: Absolute and relative average source contributions to PM10 mass in the 8-factor constrained solution.

898 Table 2: Average contribution to total reconstructed b_{ap} for the biomass burning and fossil fuel factors; in parenthesis 25th
899 and 75th percentile are reported.

900