Response to the comments of Reviewer #1

Li et al. presented a study that examined the effects of temperature on the optical properties and chemical composition of secondary organic aerosols formed from the OH photooxidation of n-dodecane. The authors found that oligomers were formed at low temperatures, and these oligomers resulted in higher RI values being measured. This paper is potentially useful to the SOA community. However, there are some important issues that the authors need to address before the manuscript can be considered for publication.

Response: We thank Anonymous Referee #1 for the review and the positive evaluation of our manuscript. We have fully considered the comments and responded to these comments below in blue text. The revisions in the manuscript are highlighted in yellow color. The response and changes are listed below.

<u>Major comments:</u>

1. Why were experiments conducted under dry conditions?

Optical properties of the SOA can be affected by many factors, in order to study the temperature affect, other factors must be kept unchanged, so the humidity of the experiments must be constant and cannot be changed. The experiments were conducted under dry conditions (RH < 5%). Because when the temperature changes, the saturated vapor pressure of water changes, that is to say, if the RH is consistent at different temperatures, the concentration of the water is not consistent; when the concentration of water is the same, the RH is different. So choosing other humidity (non-dry conditions) will introduce new problems, we can only choose dry conditions. And we have added the related statement in the main text (Page 3, Line 88-93).

2. Table 1 showed used of 43 ppb at low temperature vs. 58 ppb at high temperature. Why wasn't the same amount of n-dodecane used? Actually, the concentration of *n*-dodecane is tested with PTR-QMS, and the calibration of the PTR-QMS's response to *n*-dodecane is achieved through permeation tubes. According to our experimental design, the expected concentration of *n*-dodecane is 50 ppb, which is introducing 2 μ L liquid *n*-dodecane into 5 m³ smog chamber. As the injection volume of *n*-dodecane is 2 μ L, volume error during injection is inevitable, which will influence the concentration of *n*-dodecane in the chamber. Nevertheless, the relative small differences in *n*-dodecane concentration (43-50 ppb at low temperature and 52-58 ppb at high temperature) likely have little influence in SOA composition and optical properties. And we have added the related statement in the main text (Page 4, Line 100-105).

3. Do the authors know how the use of different temperatures will affect the loss rates of particles to the chamber walls? Is there a possibility that the observations of the SOA mass, composition and optical properties made by the authors can be explained partly by differences of particle wall loss rates at different temperatures?

We have measured the wall loss rates of particles under both room and low temperatures, and found that the wall loss rate under low temperature condition (0.0025 $\sim 0.0028 \text{ min}^{-1}$) is larger than that under room temperature condition (0.0018 $\sim 0.0020 \text{ min}^{-1}$). However, this difference in particle wall loss rate can only slightly change the SOA mass concentration, but not the particle chemical composition. Therefore, it is unlikely to change the optical properties.

Nevertheless, the difference in vapor wall loss rates may change the particle composition and optical properties. The low temperature can enhance the loss rates of higher-volatility compounds (while for lower-volatility compounds, their dominant fates are condensation so temperature may affect little on their losses), which may lead to their lower proportions in SOA particles. As their RIs are generally lower than lower-volatility compounds (Li et al., 2018), this proportion change can probably enhance SOA RI at low temperature. Therefore, the difference in wall losses and gas-particle partitioning of gas-phase products might partially contribute to the RI enhancement

under low temperature condition.

We have added the above contents in the revised paper (Page 10, Line 297-308).

Reference:

Li, K., Li, J., Wang, W., Li, J., Peng, C., Wang, D., and Ge, M.: Effects of gas-particle partitioning on refractive index and chemical composition of m-xylene secondary organic aerosol, J. Phys. Chem. A, 122, 12, 3250-3260, 10.1021/acs.jpca.7b12792, 2018.

4. In page 5 line 144, the authors state that "During this period, the optical properties of the particles tend to be stable and will not change much." This sentence is ambiguous and needs to be clarified. What optical property is the authors referring to? RI value? Or are they referring to the absorption spectra?

The optical property here refers to the extinction coefficients (α_{ext}) of the particles. During the last 1 h of the experiments, the extinction coefficients (α_{ext}) measured by the CRDS at 532 nm and PAX at 375 nm tend to be stable; at the same time, the surface mean diameter (D) of the particles tend to be stable and does not change much. When the D and α_{ext} are stable, the extinction efficiency (Q_{ext}) will be stable (as shown in the following equation). A fixed set of D and Q_{ext} results in a fixed RI value.

$$Q_{ext} = \frac{4\alpha_{ext}}{N\pi D^2}$$

We have rephrased this sentence to "During this period, the surface mean diameter and the extinction coefficients (α_{ext}) of the particles tended to be stable and will not change much" (Page 7, Line 210-211)

5. Explain the rationale behind tracking the absorption at 532 and 375 nm.

The CRDS measures the total optical extinction coefficient (α_{ext}) of the SOA, and $\alpha_{ext} = \alpha_{abs}$ (absorption coefficient) + α_{sca} (scattering coefficient). In order to calculate the real (n) and imaginary (i) part of RI values, both the α_{abs} and the α_{sca} need to be known. So the absorption at 532 nm is measured with a UV-Vis light spectrometer (Avantes 2048F). With the measured α_{ext} and α_{abs} , the α_{sca} will be calculated.

For the absorption at 375 nm, the rationale is as following: The photoacoustic extinctiometer (PAX-375, Droplet Measurement Technologies) directly measures insitu light absorption and scattering of aerosol particles at 375 nm, from which the α_{abs} and α_{sca} can be derived.

6. The authors did not use seed aerosols in this study to promote gas-to-aerosol partitioning. Hence, I expect substantial vapor wall loss in these experiments, and the extent of vapor wall loss is likely to be different at 5 °C vs. 25 °C. Is it possible that the authors are not detecting some products (due to their loss to the chamber walls) that can contribute to SOA optical properties?

We agree that the gas-particle partitioning can influence the particle composition and optical properties. As shown in our previous study (Li et al., 2017), the presence of seeds can promote the condensation of low-molecular-weight products and decrease the real part of the RI of *n*-dodecane SOA under low NO_X conditions. The experimental conditions in that study are very similar to those of room temperature condition in this study. Therefore, we can expect that similar results can be found in this study at room temperature. However, even if we perform the experiments with seeds in this study, the gas-particle partitioning can also be different at different temperature. In other words, we cannot rule out the contribution of different gas-particle partitioning at different temperatures.

As we have discussed in our response to comment #3, the difference in gas-particle partitioning of products might partially contribute to the RI enhancement under low temperature condition.

Reference:

- Li, J., Li, K., Wang, W., Wang, J., Peng, C., and Ge, M.: Optical properties of secondary organic aerosols derived from long-chain alkanes under various NO_x and seed conditions, Sci. Total Environ., 579, 1699-1705, 10.1016/j.scitotenv.2016.11.189, 2017.
- 7. In section 3.5, the authors tried to relate their results to observations made during

winter haze episodes in China. I advise the authors to be more circumspect in the extrapolation of their results to ambient observations since NOx concentrations are likely substantial during winter haze episodes in China. The authors performed a study under low-NOx conditions. Under high NOx conditions, I expect the reaction mechanism of n-dodecane OH photooxidation to be different. For example, more fragmentation will likely happen, which will result in the formation of more volatile products. If this is the case for both 5 °C and 25 °C conditions, there may not be significant differences between SOA composition and their RI values under high NOx conditions, which would imply that temperature does not play a big role in the DRF of SOA formed under areas with significant NOx concentrations, like China.

We agree that the reaction mechanism of n-dodecane under high NO_X conditions is different from that under low NO_X conditions. The temperature effects in the presence of NO_X are indeed very important and will be investigated in future studies. Therefore, we have revised this part in the manuscript "The enhancement in light extinction of SOA and oligomer compositions formed under low temperature condition might provide some possible inspiration for the regional visibility issues, especially the suburban areas. It had been reported that UV-scattering particles in the boundary layer could accelerate photochemical reactions and haze production (Sun et al., 2014). The observations above showed that the scattering property of formed SOA increased under low temperature condition, which might provide one possible reason for the rapid occurrence of haze in suburban areas in winter. According to field observations, haze occurred frequently in winter, especially in China (Cheng et al., 2016; Huang et al., 2014; Guo et al., 2014; Parrish et al., 2007). When haze occurred, it was often accompanied by high NO_X, especially in urban areas. The temperature effects under high-NO_X conditions are also important and need to be investigated in future studies." (Page 11, Line 317-324)

Minor comments:

8. Inconsistent tenses. The authors switch between using past and present tenses in some parts of the manuscript. Please correct this.

Thank you for the helpful comments. We have corrected this in the manuscript.

Page 7, Line 209: are \rightarrow were; reaches \rightarrow reached;

Page 7, Line 210: are \rightarrow were;

Page 7, Line 211: tend \rightarrow tended;

Page 7, Line 215: are \rightarrow were;

Page 7, Line 216: have \rightarrow had;

Page 8, Line 223: indicate \rightarrow indicated;

Page 8, Line 249: are \rightarrow were;

Page 9, Line 257: changes \rightarrow changed.

Response to the comments of Reviewer #2

This work by Li et al. describes measurements of optical properties in relation with chemical composition of n-dodecane SOA under low- NO_x and two different temperatures (5 °C and 25 °C). The authors found that under low temperature, the real part of the refractive index (RI) at 375 nm and 532 nm is enhanced, corresponding to substantial oligomer formation. The authors hence conclude that the enhanced oligomer formation under low temperature lead to the higher RI. The results could be relevant to low visibility in urban areas during wintertime. Overall, the manuscript is well written and demonstrates new findings regarding SOA optical properties vs. chemical compositions. But a few major concerns need to be addressed before this manuscript can be considered publishing.

Response: We thank Anonymous Referee #2 for the review and the positive evaluation of our manuscript. We have fully considered the comments and responded to these comments below in blue text. The revisions in the manuscript are highlighted in yellow color. The response and changes are listed below.

<u>Specific comments:</u>

1. Line 49 – 52. Although it is true that few studies have examined temperature effects on RI values of SOA, temperature effect studies on SOA formation and composition are not limited. There are many other studies on SOA formation and temperature effects not cited. For example, Warren et al., 2009, 43, 3548, Atmos. Environ.; Emanuelsson et al., 2013, 117, 10346, J. Phys. Chem. A.; Price et al., 2016, 50, 1216, Aerosol. Sci. Technol.; Boyd et al., 2017, 51, 7831, Environ. Sci. Technol.; Zhao et al., 2019, 3, 2549, ACS Earth Space Chem., etc.

We have revised this part in the manuscript, and the references mentioned above have been cited:

"There are also many studies investigating temperature effects on SOA formation and composition (Takekawa et al., 2003; Svendby et al., 2008; Clark et al., 2016; Lamkaddam et al., 2016; Huang et al., 2017; Qing Mu and Gerhard Lammel, 2018; Zhao et al., 2019; Boyd et al., 2017; Price et al., 2016; Emanuelsson et al., 2013; Qi et al., 2010; Warren et al., 2009); however, works on the effect of temperature on the SOA RI are limited (Kim et al., 2014)." (Page 2, Line 52-56)

References:

- Boyd, C. M., Nah, T., Xu, L., Berkemeier, T., and Ng, N. L.: Secondary organic aerosol (SOA) from nitrate radical oxidation of monoterpenes: Effects of temperature, dilution, and humidity on aerosol formation, mixing, and evaporation, Environ. Sci. Technol., 51, 7831-7841, 10.1021/acs.est.7b01460, 2017.
- Emanuelsson, E. U., Watne, A. K., Lutz, A., Ljungstrom, E., and Hallquist, M.: Influence of humidity, temperature, and radicals on the formation and thermal properties of secondary organic aerosol (SOA) from ozonolysis of beta-pinene, J. Phys. Chem. A, 117, 10346-10358, 10.1021/jp4010218, 2013.
- Huang, W., Saathoff, H., Pajunoja, A., Shen, X., Naumann, K.-H., Wagner, R., Virtanen, A., Leisner, T., and Mohr, C.: α-Pinene secondary organic aerosol at low temperature: Chemical composition and implications for particle viscosity, Atmos. Chem. Phys., 18, 2883-2898, https://doi.org/10.5194/acp-18-2883-2018, 2018.
- Kim, H., Liu, S., Russell, L. M., and Paulson, S. E.: Dependence of real refractive indices on O:C, H:C and mass fragments of secondary organic aerosol generated from ozonolysis and photooxidation of limonene and alpha-pinene, Aerosol Sci. Technol., 48, 498-507, 10.1080/02786826.2014.893278, 2014.
- Lamkaddam, H., Gratien, A., Pangui, E., Cazaunau, M., Picquet-Varrault, B., and Doussin, J.-F.: High-NOx photooxidation of n-dodecane: Temperature dependence of SOA formation, Environ. Sci. Technol., 10.1021/acs.est.6b03821, 2016.
- Price, D. J., Kacarab, M., Cocker, D. R., Purvis-Roberts, K. L., and Silva, P. J.: Effects of temperature on the formation of secondary organic aerosol from amine precursors, Aerosol Sci. Technol., 50, 1216-1226, 10.1080/02786826.2016.1236182, 2016.
- Qi, L., Nakao, S., Tang, P., and R., C. I. D.: Temperature effect on physical and chemical properties of secondary organic aerosol from m-xylene photooxidation, Atmos. Chem. Phys., 10, 3847-3854, 2010.
- Qing Mu, M. S., Mega Octaviani, Nan Ma, Aijun Ding, Hang Su,, and Gerhard Lammel, U. P., Yafang Cheng: Temperature effect on phase state and reactivity controls atmospheric multiphase chemistry and transport of PAHs, Sci. Adv., 4, 10.1126/sciadv.aap7314, 2018.
- Svendby, T. M., Lazaridis, M., and Tørseth, K.: Temperature dependent secondary organic aerosol formation from terpenes and aromatics, J. Atmos. Chem., 59, 25-46, 10.1007/s10874-007-9093-7, 2008.
- Takekawa, H., Minoura, H., and Yamazaki, S.: Temperature dependence of secondary organic aerosol formation by photo-oxidation of hydrocarbons, Atmos. Environ., 37, 3413-3424, 10.1016/s1352-2310(03)00359-5, 2003.
- Warren, B., Austin, R. L., and Cocker, D. R.: Temperature dependence of secondary organic aerosol,

Atmos. Environ., 43, 3548-3555, 10.1016/j.atmosenv.2009.04.011, 2009.

Zhao, Z., Le, C., Xu, Q., Peng, W., Jiang, H., Lin, Y.-H., Cocker, D. R., and Zhang, H.: Compositional evolution of secondary organic aerosol as temperature and relative humidity cycle in atmospherically relevant ranges, ACS Earth Space Chem., 3, 2549-2558, 10.1021/acsearthspacechem.9b00232, 2019.

2. Line 78. Although details of the chamber can be referred to previous work, important and fundamental characteristics of the chamber still needs to be provided. For example, the material and volume of the chamber.

We have added the important and fundamental characteristics of the chamber in the **2.1 Smog Chamber Experiments** part:

"The chamber consisted of two 5 m³ reactors made of fluorinated ethylene propylene (FEP) Teflon-film, which were housed in a thermally isolated enclosure. The temperature in the chamber was accurately controlled by high-power air conditioner in the range of -10 - 40 °C." (Page 3, Line 82-84)

3. Some important details of the chamber experiments are missing. Why were instruments not connected before temperature was stabilized? SOA mass concentration in each experiment should be reported (in Table 1 or in the main text). How long were filter samples collected and at what flow rate? After the filters were dissolved in filter, were the solutions sonicated? If not, how good were the extraction recoveries?

Actually, the instruments were connected at room temperature condition, when the temperature dropped to 5 °C and stabilized, the data measured by the instruments would to counted as valid data. And we have added this explanation in the main text (Page 3, Line 97-99).

We have added SOA mass concentration in Table 1 (Page 17, Line 605):

Table 1. The initial conditions of the sinog-chamber experiment							
HC (ppb)				Temperature (°C)	Mass ^b (µg/m ³)		
58	1.03	<1	<5	25	155		
52	1.03	<1	<5	25	135		
	HC (ppb) 58	HC H ₂ O ₂ (ppb) (ppm) 58 1.03	HC H ₂ O ₂ NOx (ppb) (ppm) (ppb) 58 1.03 <1	HC H ₂ O ₂ NO _x RH (ppb) (ppm) (ppb) (%) 58 1.03 <1	HC H2O2 NOx RH Temperature (ppb) (ppm) (ppb) (%) (°C) 58 1.03 <1		

Table 1. The initial conditions of the smog-chamber experiments.

Dod-L-1	43	1.07	<1	<5	5	128
Dod-L-2	50	1.09	<1	<5	5	133

Each filter sample was collected for 30 min at 10 L/min flow rate. Then the filters were put into 5 mL methanol (99.9%, Fisher Chemical) and sonicated for 30 min. We have added this part in the manuscript (Page 4, Line 119-120).

4. How did the extinction coefficient and extinction efficiency evolve over the course of an experiment?

With the progress of the reaction, the surface mean diameter of the particles gradually increased, and the extinction coefficient and extinction efficiency of the particles increased as well. We chose two sets of experiments under two temperature conditions as examples. The figure below has been added in the Supporting Information of the manuscript. (Supporting Information, Page 7, Figure S4.)

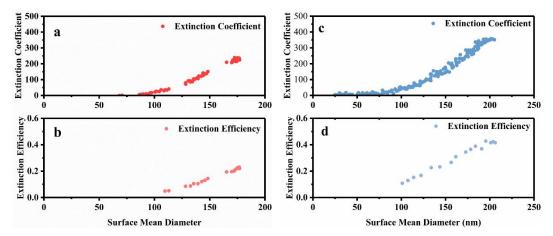


Figure S4. Evolution of optical parameters at 532 nm. Extinction coefficient (a) and efficiency (b) at room temperature; extinction coefficient (c) and efficiency (d) at low temperature.

5. Line 122. It is unclear how was the RI value uncertainty estimated from the various uncertainties. An equation needs to be provided.

We have added the estimating method and equations in the Supporting Information: (Supporting Information, Page 2, Line 28-43)

"The RI values are obtained with extinction efficiency (Qext) and surface mean

diameter (D_s). The uncertainty of the D_s measured by SMPS is $\pm 1\%$. The uncertainty of the extinction efficiency is calculated as the following:

The extinction efficiency can be expressed as:

$$Q_{\text{ext}} = \alpha_{\text{ext}} / (\frac{1}{4} N \pi D^2)$$

The three variables α_{ext} , N, D are independent and the uncertainty of them were 3%, 10% and 1%, respectively. Considering the propagation of uncertainty, we can obtain the variance formula:

$$\sigma_{\rm Q}^2 = \sigma_{\alpha}^2 \ (\frac{\partial Q}{\partial \alpha})^2 + \sigma_{\rm N}^2 \ (\frac{\partial Q}{\partial \rm N})^2 + \sigma_{\rm D}^2 \ (\frac{\partial Q}{\partial \rm D})^2$$

Then the uncertainty of Qext can be calculated as:

$$\frac{\sigma_{\rm Q}}{\rm Q} = \sqrt{\left(\frac{\sigma_{\rm \alpha}}{\rm \alpha}\right)^2 + \left(\frac{\sigma_{\rm N}}{\rm N}\right)^2 + 4\left(\frac{\sigma_{\rm D}}{\rm D}\right)^2} = \sqrt{0.03^2 + 0.1^2 + 4 \times 0.01^2}$$
$$= \sqrt{0.0113} = 10.6\%$$

The measured extinction efficiency ($Q_{ext,mea}$) is compared to calculated extinction efficiency ($Q_{ext,cal}$). The best-fit RI value is determined by minimizing the following reduced merit function (χ_r):

$$\chi_{r} = \frac{1}{N} \sum_{i=1}^{N} (Q_{ext,mea} - Q_{ext,cal}(n,k))_{i}^{2}$$

The uncertainty of the retrieval method is ± 0.002 , and the uncertainty of the measured extinction efficiency is $\pm 10.6\%$, resulting in the final uncertainty of the retrieved RI value to be about 0.02-0.03."

6. Line 123 – 124. The description of the RI prediction is not sufficient. A little more background should be provided. It appears molecular formulas are needed as input for the prediction. More details of the input are needed. Also, how the predictions will be used in this study were not mentioned.

We have added the following section to the main text (Page 6, Line 166-182): "2.3.2 Calculation Method of the QSPR Based on the Molecular Formula

Using the molecular formula obtained from ESI-TOF-MS, we calculated the RI

values of the products in SOA with the quantitative structure–property relationship (QSPR) method (Redmond and Thompson, 2011). The QSPR can be expressed with Equation (9):

$$RI_{predicted} = 0.031717(\mu) + 0.0006087(\alpha) - 3.0227\left(\frac{\rho_m}{M}\right) + 1.38709$$
(9)

where μ is the unsaturation of the molecular formula, α is the polarizability of the molecular formula, ρ_m is the mass density (g/cm³), and *M* is the molar mass (g/mol).

The mass density of the compound is estimated by AIM model, the detailes of which can be referred to Girolami (1994).

 μ is calculated through the conventional approach, which is used in many organic chemistry texts, Equation (10)

$$\mu = (\#C + 1) - 0.5(\#H - \#N) \tag{10}$$

where #C,#H, and #N are the number of the C, H, and N respectively.

 α is calculated based on the molecular formula of the compound, it can be expressed by Equation (11):

$$\alpha = 1.51(\#C) + 0.17(\#H) + 0.57(\#O) + 1.05(\#N) + 2.99(\#S) + 2.48(\#P) + 0.22(\#F) + 2.16(\#Cl) + 3.29(\#Br) + 5.45(\#I) + 0.32$$
(11)

where # is the number of the atoms of each element in the molecular formula.

The calculated RI values of products were used to link chemical composition and optical properties, and to explain the obseved RI differences at different temperatures in Sect. 3.4."

Reference:

7. Line 150. In prior description (Line 122), the authors claimed that the uncertainty for RI values are 0.02 - 0.03. Here, the authors argue that change by 0.02 - 0.03 is a "substantial enhance effect" between the two temperatures, despite this change is on par with the uncertainty. A better case needs to be demonstrated regarding the enhancement of RI values at low temperature. In particular, this argument is a main result in this work.

Redmond, H., and Thompson, J. E.: Evaluation of a quantitative structure-property relationship (QSPR) for predicting mid-visible refractive index of secondary organic aerosol (SOA), Phys. Chem. Chem. Phys. : PCCP, 13, 6872-6882, 10.1039/c0cp02270e, 2011.

Girolami, G. S.: A simple" back of the envelope" method for estimating the densities and molecular volumes of liquids and solids, J. Chem. Edu., 71, 11, 962, 1994.

In order to more intuitively display the RI variation tendency of SOA generated under two temperature conditions, we have added Figure 2 in the main text. As shown in Figure 2, the RI under low temperature conditions is enhanced whether at 532 nm or 375 nm. The enhancement of RI values under low temperature condition is still valid with experiment uncertainty. We have changed "a substantial enhance effect" into "an enhance effect" in the revised manuscript. (Page 8, Line 219)

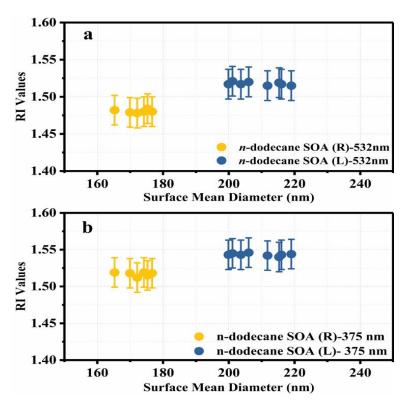


Figure 2. Variation tendency of the RI values as a function of surface mean diameter for SOA produced under room and low temperature conditions at (a) 532 nm and (b) 375 nm.

8. Section 3.3. and Figure 2. From the chemical composition measured by ESITOFMS, the authors claimed that the largely different mass spectra are observed under lower temperature, due to enhanced oligomerization. It is unclear, however, how the changed temperature affected the chemical mechanisms that lead to different products. The discussion of higher oligomer formation under low temperature might be reasonable. But the difference in the monomer range does not make sense. The monomer products should be from the various RO_2 chemistry regardless of the temperature. They should follow the same pattern. But in Figure 2, the monomer range in the two mass spectra show very different results. I wonder if the mass spectra shown here are reproducible? Two experiments were carried out for each condition. Do they show consistent mass spectra? Clark et al (2016, 50, 5564, Environ. Sci. Technol.) showed isoprene SOA ESI-TOFMS mass spectra under different temperatures. Similar to this work here, very different results were observed in the oligomer range, but not necessarily for monomers. Better clear discussions are need here, rather than speculations.

We have updated Figure 2b (now is Figure 3b) by removing impurity interferences at ~200 and 300 amu:

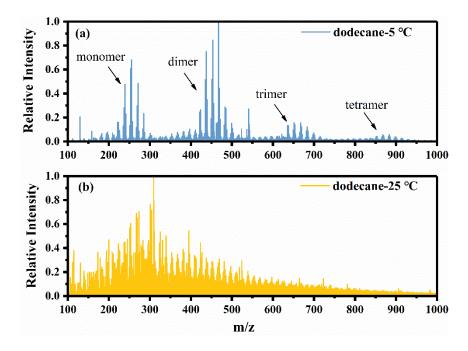
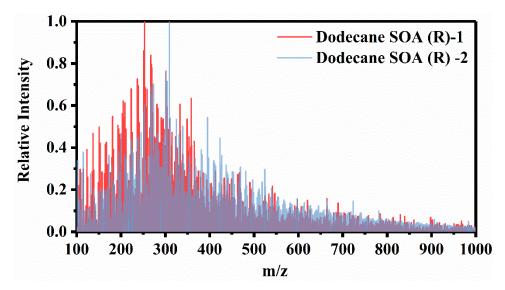


Figure 3. Mass spectra of *n*-dodecane SOA obtained by ESI-TOF-MS in positive ion mode. (a) low temperature condition; (b) room temperature condition.

The mass spectra under the same temperature condition are reproducible, see below the mass spectra of two experiments under room temperature. Though there are some slight differences in peak intensities, these two mass spectra are generally similar.



As the reviewer said, Clark et al. (2016) found that the SOA formed from isoprene had similar mass spectra at the monomer range at different temperatures. However, the precursors are very different (isoprene vs *n*-dodecane), which have very different oxidation pathways and partitioning processes. Hence, it is possible that we observe different results in mass spectra at monomer range.

As we have mentioned in the manuscript, the different oxidation degree and partitioning process have made the difference in monomer range. At room temperature, more OH oxidation steps in the gas phase can lead to the formation of some fragmentation products that may not be observed in the low temperature. For example, we have observed the high intensity of M/Z 195 $C_{10}H_{20}O_2$ (Relative Intensity 0.39), M/Z 211 $C_{10}H_{20}O_3$ (Relative Intensity 0.13) at room temperature but very low intensity (0.012, 0.041, respectively) at low temperature, and M/Z 195 and 211 are likely products from gas-phase OH oxidation. This process may be different from isoprene (Clark et al., 2016), because for *n*-dodecane the carbon number is high and volatility is low and fewer oxidation steps are needed before partitioning, while for isoprene there is only 5 carbon atoms so more oxidation steps are needed before partitioning no matter at room or low temperature.

Overall, the differences in monomer range still make sense.

We have added related statement in the main text. (Page 9, Line 262-273)

Reference:

Clark, C. H., Kacarab, M., Nakao, S., Asa-Awuku, A., Sato, K., and Cocker, D. R., 3rd: Temperature effects on secondary organic aerosol (SOA) from the dark ozonolysis and photo-oxidation of isoprene, Environmental Science & Technology, 10.1021/acs.est.5b05524, 2016.

9. Section 3.4. It appears to me that the validation of the authors' argument using calculation method is important to extend further. Since it is a short manuscript, I suggest the authors include the results based on the prediction in the main text.

Thanks for this suggestion. We have moved the calculation method to the main text, and they are highlighted in yellow in the "2.3 Calculation Method of RI Values" Section. (Page 4-6, Line 125-189). Figure S3, Figure S5, and Table S2 are also moved to the main text, and they have updated as Figure 4, Figure 6, and Table 2.

Technical comments:

1. Line 29. Change "heavy" to "heavily".

Page 1, Line 32: This has been corrected.

2. Line 62 – 68. When referring to prior temperature effect studies, it would be helpful to mention at what temperatures were those results observed.

Page 2, Line 67: The temperature ranges have been added.

3. In Figure 4, the y-axis between a and c, as well as b and d, are different. I suggest the authors use the same y-aixs ranges for easier comparison.

Thanks for this suggestion. This has been corrected, and Figure 4 has been updated as Figure 7 in the main text.

Temperature Effects on Optical Properties and Chemical Composition of Secondary Organic Aerosol Derived from *n*-Dodecane

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Abstract. Environmental temperature plays a vital role in controlling chemical transformations that lead to the formation of

- 20 secondary organic aerosol (SOA), and ultimately impact composition and optical properties of the aerosol particles. In this study, we investigated optical properties of *n*-dodecane secondary organic aerosol under two temperature conditions: 5 °C and 25 °C. It was shown that low temperature could enhance the real part of refractive index (RI) of the SOA at the wavelengths of 532 nm and 375 nm. Mass spectrometry analysis revealed that molecular composition of *n*-dodecane SOA was significantly modified by temperature: a large amount of oligomers were formed under low temperature condition,
- 25 which leaded to higher RI values. These findings will help improve our understanding of the chemical composition and optical properties of SOA under different temperature conditions, and provide one possible explanation of the low visibility in suburban areas during winter.

1 Introduction

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Organic aerosol, especially secondary organic aerosol (SOA), plays a vital role in air quality, climate change, and human health (Kanakidou et al., 2005; Poschl, 2005; Mellouki et al., 2015; Poschl and Shiraiwa, 2015; von Schneidemesser et al., 2015; Shrivastava et al., 2017). SOA accounts for a high proportion of atmospheric particulate matter around the world, especially in the heavily polluted areas (Liu et al., 2017; Sun et al., 2014; Huang et al., 2014). Due to the variety of precursors and oxidation pathways, the composition of SOA is very complicated and variable (Lu, 2018; von Schneidemesser et al., 2015; Poschl and Shiraiwa, 2015; Hallquist et al., 2009; George et al., 2015), and the optical 35 properties of SOA also exhibit different characteristics (Shrivastava et al., 2017; Zhang et al., 2015; Moise et al., 2015; Laskin et al., 2015). The complex refractive index (RI), m= n + ki (n is the real part, and k is the imaginary part; they express the extent of scattering and absorbing, respectively), is the only intrinsic optical property of a particle. RI is controlled by the chemical composition and physical characteristics (e.g., morphology and shape) of a particle (Moise et al., 2015). Quantifying the RI of aerosol particles is highly important to evaluate their optical properties, and further estimate their impacts on atmospheric visibility and Earth's radiative balance.

Aerosol physicochemical properties are strongly dependent on the atmospheric conditions, e.g., relative humidity (Ervens et al., 2011; Sun et al., 2014), temperature (Wang et al., 2017), and oxidizing conditions (oxidant type, e.g., NO₃, OH, O₃; oxidation concentrations, e.g., photochemical age) (Cheng et al., 2016; Shrivastava et al., 2017; George et al., 2015; Kanakidou et al., 2005). Therefore, it is important to study the SOA formation and optical properties under varying atmospheric conditions to simulate the processes in the real atmosphere. There have been a number of smog chamber experiments on the effect of seed particles (Huang et al., 2017; Denjean et al., 2014; Song et al., 2013; Lee et al., 2013; Li et al., 2018; Li et al., 2017a; Trainic et al., 2011), oxidant type (e.g., NO₃ (Peng et al., 2018; Lu et al., 2011), OH (Liu et al., 2015; Lin et al., 2015; Li et al., 2014; Nakayama et al., 2013; Cappa et al., 2011), and O₃ (Peng et al., 2018; Kim et al., 2014; Kim and Paulson, 2013)), oxidation concentrations (e.g., photochemical age (Zhong and Jang, 2014; Kim et al., 2014; Lambe et al., 2012; Lambe et al., 2013)), and relative humidity (RH) (Titos et al., 2016; McNeill, 2015; Denjean et al., 2015; Li et al., 2012; Lambe et al., 2013)), and relative humidity (RH) (Titos et al., 2016; McNeill, 2015; Denjean et al., 2017b; Song et al., 2017b; Nong et al., 201

- et al., 2015; Li et al., 2017b; Sareen et al., 2017; Ye et al., 2016; Michel Flores et al., 2012) on SOA formation and the RI values of SOA derived from both biogenic and anthropogenic volatile organic compounds (VOCs). There are also many studies investigating temperature effects on SOA formation and composition (Takekawa et al., 2003; Svendby et al., 2008; Clark et al., 2016; Lamkaddam et al., 2016; Huang et al., 2018; Qing Mu and Gerhard Lammel, 2018; Zhao et al., 2019;
- 55 Boyd et al., 2017; Price et al., 2016; Emanuelsson et al., 2013; Warren et al., 2009; Qi et al., 2010); however, works on the effect of temperature on the SOA RI are limited (Kim et al., 2014). Field studies have shown that temperature is an important factor affecting rate constants of the oxidation process, vapor pressure of products, and SOA formation process and yields (Atkinson and Arey, 2003; Wang et al., 2017; Roy and Choi, 2017; Ding et al., 2017; Cui et al., 2016). Therefore, investigating temperature dependence is important to our better understanding of the formation, physical, and chemical properties of SOA under tropospheric conditions.
 - Long-chain alkanes, an important class of intermediate-volatility organic compounds (IVOCs) (Zhao et al., 2014) and a large fraction of diesel fuel and its exhaust (Gentner et al., 2012; Gentner et al., 2017), are important potential contributor of SOA (Presto et al., 2009; Zhao et al., 2016). Several previous studies have reported the formation of SOA derived from long chain alkanes under various conditions, including SOA compositions (Fahnestock et al., 2015; Hunter et al., 2014; Lim and
- 65 Ziemann, 2005), SOA yields (Loza et al., 2014; Lim and Ziemann, 2009a), and the chemical mechanisms (Yee et al., 2012, 2013; Aimanant and Ziemann, 2013; Lim and Ziemann, 2009b). Recently, Lamkaddam et al. (2016) reported the temperature dependence (10 °C, 20 °C, and 31.5 °C) of SOA formation from high NO_x photo-oxidaton of *n*-dodecane, and found that temperature did not significantly influence SOA yield. They attributed it to two possible reasons: the changes of

reaction rate constants lead to different SOA composition; or the formed SOA are mainly non-volatile compounds that they

- are not sensitive to temperature. Li et al. (2017a) reported the optical properties of SOA from *n*-dodecane, *n*-pentadecane, and *n*-heptadecane under various oxidation conditions under room temperature. However, knowledge about the effect of temperature on the chemical composition and optical properties of *n*-dodecane SOA in the absence of NO_x is still lacking, which limits our understanding of the role of SOA in visibility and radiative balance under different temperatures (e.g., in winter and summer).
- In the present study, we determined the temperature effects on chemical composition and optical properties of SOA generated in a smog chamber during photo-oxidation of *n*-dodecane under low- NO_x condition. The results here will improve our understanding of the role of temperature in SOA chemical compositions and optical properties, and further the influence on air quality and radiative forcing.

2 Materials and methods

80 2.1 Smog Chamber Experiments

The experiments were performed in a dual-reactor smog chamber, the details of which were given previously (Wang et al., 2015). Briefly, the chamber consisted of two 5 m³ reactors made of fluorinated ethylene propylene (FEP) Teflon-film, which were housed in a thermally isolated enclosure. The temperature in the chamber was accurately controlled by high-power air conditioner in the range of -10 - 40 °C. Multiple light sources were used in the chamber, with center wavelength of 365 nm

- 85 (GE, F40BL), 340 nm (Q-lab, UVA-340), and 254 nm (PHILIPS, G36 T8). The RH and temperature in the chamber were continuously monitored and controlled during the whole experiments. The experiments were conducted under < 5% RH and under two temperatures: 25 °C (room temperature condition, R) and 5 °C (low temperature condition, L). The temperature fluctuation was ± 0.5 °C for either condition. As optical properties of the SOA could be affected by many factors, in order to study the temperature affect, other factors must be kept unchanged, so the humidity of the experiments must be constant and
- 90 cannot be changed. Here the experiments were conducted under dry conditions. Because when the temperature changed, the saturated vapor pressure of water changed: if the RH was consistent at different temperatures, the concentration of the water was not consistent; when the concentration of water was the same, the RH was different. So choosing other humidity (non-dry conditions) would introduce new problems, we could only choose dry conditions (RH < 5%). The *n*-dodecane (≥99%, Sigma-Aldrich) was photo-oxidized under low-NOx condition, with hydrogen peroxide (30% wt/wt, Beijing Chemical Chemic
- 95 Works) as the OH precursor. *n*-Dodecane was added into the chamber first, followed by adding hydrogen peroxide. After that, wind turbine was turned on for 20 min to make sure that the materials in the chamber were well mixed.

When all the substances were added in the chamber, the chamber was set to the desired temperature. The instruments were connected at room temperature condition, when the temperature dropped to 5 °C and stabilized, the data measured by the instruments would be counted as valid data. The lights in the chamber were then turned on, and the photo-oxidation

100 reaction started. The initial conditions for these experiments were listed in Table 1. According to our experimental design,

the expected concentration of *n*-dodecane was 50 ppb, which was introducing 2 μ L liquid *n*-dodecane into 5 m³ smog chamber. As the injection volume of *n*-dodecane was 2 μ L, volume error during injection was inevitable, which would influence the concentration of *n*-dodecane in the chamber. Nevertheless, the relative small differences in *n*-dodecane concentration (43-50 ppb at low temperature and 52-58 ppb at high temperature) likely had little influence in SOA composition and optical properties.

2.2 Measurements

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All instruments were located within one meter of the chamber, and all the connection tubes were wrapped by insulation cotton to minimize the influence of room temperature. The concentration of NOx and formed O_3 in the smog chamber were monitored by the gas analyzers (Teledyne Advanced Pollution Instrumentation, Model T400 and Model T200UP, respectively). The concentration of *n*-dodecane was monitored by a proton transfer reaction quadrupole mass spectrometry (PTR-OMS 500, Ionicon), calibration of PTR-OMS's response to *n*-dodecane was achieved through permeation tubes. NO⁺

ion source of PTR-QMS was used when detecting the *n*-dodecane (Koss et al., 2016; Shi et al., 2019; Paulsen et al., 2005). The particle size distribution and density were detected by a scanning mobility particle sizer (SMPS, TSI) and a

centrifugal particle mass analyzer (CPMA, Cambustion). A custom-made cavity ring-down spectrometer (CRDS) (Wang et
al., 2012) was applied to monitor the optical property of the formed particles at 532 nm. A photoacoustic extinctiometer
(PAX-375, Droplet Measurement Technologies) was used to measure the scattering, absorption, and extinction coefficients of formed SOA at 375 nm.

The formed aerosol particles were collected on the PTFE membrane with a pore size of 200 nm (0.2 μ m, 47 mm, Merckmillipore FELP). Each filter sample was collected for 30 min at 10 L/min flow rate. Then the filters were put into 5

120 mL methanol (99.9%, Fisher Chemical) and sonicated for 30 min. The dissolved solution was analyzed with a UV–Vis light spectrometer (Avantes 2048F), which was used to detect the absorbing property (to derive the imaginary part of RI, k) at 532 nm. The solution was also analyzed with electrospray ionization time-of-flight mass spectrometry (ESI-TOF-MS, Bruker, Impact HD) to obtain the chemical composition of the formed SOA. Positive ion mode was used for the ESI-TOF-MS. The absolute mass error was below 3 ppm, and the typical mass resolving power is >3000000 at m/z 200.

125 2.3 Calculation Method of RI Values

2.3.1 Calculation Method of RI Values based on the CRDS

The RI values of the particles formed in the smog chamber was estimated based on both the extinction and scattering coefficients and Mie-Lorenz theory (Bohren, 1983). The details of the calculation method of RI values are also reported in our previous publications (Wang et al., 2012; Phillips and Smith, 2014; Li et al., 2017a; Li et al., 2017b; Li et al., 2018; Peng

130 et al., 2018).

The extinction coefficients (α_{ext}) of the particles with CRDS can be calculated with Equation (1):

where *L* is the distance of the two mirrors in the cavity, *l* is the length of the cavity that filled with aerosol particles, *c* is the speed of the light, τ_0 is the ring down time of the CRDS when it is filled with zero air, and τ is the ting down time of the CRDS when it is filled with aerosol particles

For the homogeneous spherical particles, the α_{ext} can be calculated with Equation (2):

$$\alpha_{ext} = N\sigma_{ext} = \frac{1}{4}N\pi D^2 Q_{ext} \tag{2}$$

where N is the concentration of the spherical particles, σ_{ext} is the extinction cross section, D is the particle diameter, and Q_{ext} is the extinction efficiency.

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 Q_{ext} is the ratio of Beer's law extinction cross section to the geometric area of the spherical particles, it is dimensionless, and can be expressed with Equation (3):

$$Q_{ext} = \frac{4\alpha_{ext}}{N\pi D^2} \tag{3}$$

For the polydisperse SOA particles formed in the smog chamber, they follow log-normal distribution and the geometric standard deviation is always smaller than 1.5, the Q_{ext} with a surface mean diameter D_s can be expressed with Equation (4)

(with the assumption that the aerosol particles during each size bin are homogeneous spherical particles) (Nakayama et al., 2010):

$$Q_{ext}(D_s) = \frac{a_{ext}}{\int N(D_p) \frac{1}{4} \pi D_p^2 dD_p}$$
(4)

where D_p is the geometrical diameter of the particle, dD_p is the size bin of the particles, and $N(D_p)$ is the number concentration of the particles in dD_p with D_p per unit volume.

For the particles with D_p in each size bin dD_p , the surface area $S(D_p)$ can be expressed by Equation (5): $S(D_p) = N(D_p)\pi D_p^2$ (5)

So the Equation (4) can also be expressed with Eauqtion (6):

$$Q_{ext}(D_s) = \frac{\alpha_{ext}}{\int_{\frac{1}{4}}^{1} S(D_p) dD_p} = \frac{4\alpha_{ext}}{S_{tot}}$$
(6)

where S_{tot} is the total surface area of the particles, and the values can be obtained with SMPS.

While the extinction efficiency can also be calculated with Mie-Lorenz theory, Equation (7):

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$$Q_{ext}(D_s) = \int f(D_p) Q_{ext}(D_p) dD_p$$

where $f(D_p)$ is the normalized surface area weighted size distribution function.

Then the measured extinction efficiency ($Q_{ext,mea}$) is compared to calculated extinction efficiency ($Q_{ext,cal}$). And the bestfit RI value is determined by minimizing the following reduced merit function (χ_r), Equation (8):

$$\chi_r = \frac{1}{N} \sum_{i=1}^{N} (Q_{ext,mea} - Q_{ext,cal}(n,k))_i^2$$

where N is the number of diameters to be calculated.

The uncertainties of the particle concentration and surface mean diameter measured by SMPS are $\pm 10\%$ and $\pm 1\%$ respectively. The uncertainty of the retrieval method is ± 0.002 , and the uncertainty of the measured extinction coefficient

(8)

(7)

with CRDS is $\pm 3\%$, resulting in the final uncertainty of the retrieved RI value to be about 0.02-0.03. And the corresponding

165 equation for the RI value uncertainty can be referred to the Supporting Information.

2.3.2 Calculation Method of RI Values based on QSPR

Using the molecular formula obtained from ESI-TOF-MS, we calculated the RI values of the products in SOA with the quantitative structure–property relationship (QSPR) method (Redmond and Thompson, 2011). The QSPR can be expressed with Equation (9):

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$$RI_{predicted} = 0.031717(\mu) + 0.0006087(\alpha) - 3.0227\left(\frac{\rho_m}{M}\right) + 1.38709$$
 (9)

where μ is the unsaturation of the molecular formula, α is the polarizability of the molecular formula, ρ_m is the mass density (g/cm³), and *M* is the molar mass (g/mol).

The mass density of the compound is estimated by AIM model, the detailes of which can be referred to Girolami (1994).

 μ is calculated through the conventional approach, which is used in many organic chemistry texts, Equation (10)

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$$\mu = (\#C + 1) - 0.5(\#H - \#N)$$
(10)

where #C,#H, and #N are the number of the C, H, and N respectively.

α

 α is calculated based on the molecular formula of the compound, it can be expressed by Equation (11):

$$= 1.51(\#C) + 0.17(\#H) + 0.57(\#O) + 1.05(\#N) + 2.99(\#S) + 2.48(\#P) + 0.22(\#F) + 2.16(\#Cl) + 0.12(\#F) + 0.12(\#$$

$$3.29(\#Br) + 5.45(\#I) + 0.32 \tag{11}$$

180 where *#* is the number of the atoms of each element in the molecular formula.

The calculated RI values of products were used to link chemical composition and optical properties, and to explain the observed RI differences at different temperatures in Sect. 3.4.

2.3.3 Calculation Method of Average Elemental Composition and Ratios

Due to the complexity of chemical composition in secondary organic aerosol particles, it is common to express the bulk composition as averaged elemental ratios. The average elemental composition and ratios (C, H, O, H/C, O/C) can be estimated from the identified molecular formulas (Ervens et al., 2011):

$$\langle \mathbf{Y} \rangle = \frac{\sum_{i} x_{i} Y_{i}}{\sum_{i} x_{i}}$$
(12)

$$\langle Y/Z \rangle = \frac{\sum_{i} x_{i} Y_{i}}{\sum_{i} x_{i} Z_{i}}$$
(13)

where Y=C, H, O, Y/Z=H/C, O/C. x_i is the peak abundance of elemental composition.

190 2.4 Impact of Temperature on Direct Radiative Forcing

The simple forcing efficiency (SFE) is used to dertermine the relative importance of optical properties of the aerosol to direct radiative forcing at the Earth's surface (Bond and Bergstrom, 2006):

$$SFE = \frac{s_0}{4} \tau_{atm}^2 (1 - F_c) [2(1 - a_s)^2 \frac{Q_{bs}C}{M} - 4a_s \frac{Q_aC}{M}]$$
(14)

where S_0 is the solar radiation, τ_{atm} is the transmission of the atmosphere, F_c is the cloud fraction, a_s is the surface albedo, Q_{bs} and Q_a are the backscattering and absorption efficiency of the aerosol particles, M is the aerosol mass, and C is the cross section of the aerosol.

The SOA derived from *n*-dodecane have negligible absorption at the wavelengths of 532 nm and 375 nm under the two temperature conditions, so the value of Q_a is zero, and the impact of temperature on the direct radiative forcing (DRF) can be expressed with Eq. (15), and this will be discussed Section. 3.5.

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$$DRF \ ratio = \frac{SFE_{low temperature}}{SFE_{normal temperature}} = \frac{Q_{bs,low tem.}}{Q_{bs,normal tem.}}$$
 (15)

3. Results and Discussion

3.1 Photo-oxidation Experiments

The profiles of the *n*-dodecane photo-oxidation experiments at different temperatures were shown in Figure S1. As similar amount of *n*-dodecane and oxidant were added to the smog chamber and the reaction rate was lower under low temperature, 205 the reaction time of low temperature condition was significantly longer than the room temperature condition: 4 to 5 h for room temperature, and 8 to 9 h for low temperature. Under room temperature condition, the total surface concentration and mass reached a maximum after 3 h; while for low temperature condition, this time was 6 h.

3.2 Effect of Temperature on RI

Optical properties of the formed particles were analyzed after the mass concentration of the aerosol reached the maximum,

- 210 and the last 1 h data were used. During this period, the surface mean diameter and the extinction coefficients (α_{ext}) of the particles tended to be stable and would not change much. The extinction coefficients and extinction efficiency of the particles as a function of the surface mean diameter were shown in Figure S4. The SOA derived from *n*-dodecane had no significant absorption both at 532 nm and 375 nm, similar to our previous study (Li et al., 2017a), and the imaginary part of RI will not be discussed here. The real part of RI obtained in this study was shown in Figure 1. As shown in Figure 1, the RI
- 215 values at 532 nm under room temperature were similar to our previous study (Li et al., 2017a), while the two temperature conditions had significantly different ranges for RI: 1.472-1.486 (25 °C) and 1.502-1.526 (5 °C) at 532 nm; 1.51-1.53 (25 °C) and 1.532-1.56 (5 °C) at 375 nm. And the details of the variation tendency of the RI values as a function of surface mean diameter for SOA produced under room and low temperature conditions were shown in Figure 2. The various RI values at

different temperature indicated that lower reaction temperature (from 25 °C to 5 °C) had an enhance effect (~0.03 at 532 nm,

220 ~0.02 at 375 nm) on the RI of *n*-dodecane SOA. The mass spectrometry analysis below (Section 3.3) was applied to obtain chemical composition information and explain the phenomenon above.

3.3 Temperature Effect on Chemical Composition and Reaction Mechanism

The higher RI values under low temperature indicated that the temperature might change the chemical composition of SOA by changing the reaction types or shifting the balance of different pathways. The mass spectra (MS) of *n*-dodecane SOA 225 obtained by ESI-TOF-MS in positive ion mode were shown in Figure 3, which provided the molecular insight into the chemical changes under different temperature conditions. We identified about 260 individual masses for SOA under two temperature conditions, and the details were shown in Figure S2 and Table S1. The spectrum under low temperature was significantly different with the spectrum under room temperature, with large amount of ions corresponding to monomer, dimer, trimer, and tetramer. This suggested that oligomerization might play a dominant role under low temperature condition. 230 The molecular composition of *n*-dodecane SOA was significantly modified by temperature conditions, with the averaging SOA formula changing from $C_{14.98}H_{26.47}O_{5.53}$ (R) to $C_{21.25}H_{40.44}O_{7.43}$ (L). The average carbon number increased from 14.98 at room temperature to 21.25 at low temperature, indicating that the SOA molecules were larger at low temperature. The average O/C and H/C ratios at room temperature condition were 0.37 and 1.72, respectively, while at low temperature condition the ratios were 0.35 and 1.90, respectively. The products formed under low temperature tended to 235 have higher H/C ratio and lower O/C ratio compared with the products under room temperature. The details of the O/C and H/C ratios of the products formed at different temperature conditions can be referred to Figure 5a. The phenomenon above might be due to the presence of oligomers.

For the oligomers in the mass spectrum, one possible explanation was that low temperature condition promoted gasparticle particioning and changed the particle phase reaction. The gas-phase OH oxidation could reduce H/C and increase

- O/C ratios (Li et al., 2019; Lambe et al., 2015; Li et al., 2018), while particle-phase oligomerization almost won't change O/C and H/C ratios (Charron et al., 2019). Under room temperature condition, more gas-phase oxidation steps were needed to produce the less volatile products to condense into particle-phase (because of the high temperature, i.e., high saturation vapor pressure). Hence, products formed under room temperature had lower H/C and higher O/C ratios. In contrast, the more volatile (i.e., less oxidized) products were readily to condense into particle phase under low temperature, and then undergo
- 245 particle phase reactions, e.g., oligomerization, leading to the formation of products with higher H/C and lower O/C ratios. This phenomenon was consistent with Kim et al. (2014), they studied the dependence of real part of RI on O/C and H/C ratios of SOA derived from limonene and α-pinene, and found that the higher percentage of less oxygenated semivolatile substances were responsible for the higher RI values.

For particle-phase reactions, there were mainly two reaction pathways (Fahnestock et al., 2015; Yee et al., 2012, 2013):
intramolecular cyclization of multifunctional hydroperoxides (form furan derivatives); intermolecular oligomerization of multifunctional hydroperoxides (form peroxyhemiacetal, PHA). The two pathways were competitive during

the particle phase reaction process. According to the mass spectra analysis, we speculated that the low temperature promoted the progress of the oligomerization reaction, and made it the primary pathway in the particle phase, as shown in Figure 4. As discussed above, the products with higher volatility and lower oxidation state would condense on the particle phase under

low temperature condition and then participate in the particle phase reaction, which would further promote the

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Another possible explanation was that the pathway of the gas phase reaction changed under low temperature condition comparing with room temperature condition. Combining existing mass spectrometry information, we speculated that under low temperature condition the oligomers might also be formed by gas phase radical oligomerization and then rapidly deposited into the particle phase. However, these are speculations based on the existing analysis results. The specific reaction mechanism under low temperature condition needs further investigation.

Clark et al. (2016) found that the SOA formed from isoprene had similar mass spectra at the monomer range at different temperatures. However, the precursors are very different (isoprene vs *n*-dodecane), which have very different oxidation pathways and partitioning processes. Hence, it is possible that we observe different results in mass spectra at monomer range.

As we have mentioned above, the different oxidation degree and partitioning process have made the difference in monomer range. At room temperature, more OH oxidation steps in the gas phase can lead to the formation of some fragmentation products that may not be observed in the low temperature. For example, we have observed the high intensity of M/Z 195 C₁₀H₂₀O₂ (relative intensity 0.39), M/Z 211 C₁₀H₂₀O₃ (relative intensity 0.13) at room temperature but very low intensity (0.012, 0.041, respectively) at low temperature, and M/Z 195 and 211 are likely products from gas-phase OH oxidation. This process may be different from isoprene (Clark et al., 2016), because for *n*-dodecane the carbon number is high and volatility is low and fewer oxidation steps are needed before partitioning, while for isoprene there is only 5 carbon atoms so more oxidation steps are needed before partitioning no matter at room or low temperature. Overall, the differences in monomer

range still make sense.

oligomerization reaction.

3.4 Relationship between RIs and Chemical Composition of SOA

275 Refractive index of aerosol particles is fundamentally the results of a combination of particle chemical compositions and internal mixing. SOA particles formed in the smog chamber are treated as homogenous mixtures, and the RI values of which can be expressed as (Redmond and Thompson, 2011):

$$RI = \sum x_i RI_i \tag{16}$$

where x_i is the fraction of the i_{th} component, and RI_i is the refractive index of component i.

280 The RI values were calculated for the identified products (with high intensity) under room and low temperature conditions, and the details were shown in Figure 5b,c and Table S1. It could be clearly seen that the RI values of the products at room temperature condition were mainly in the range of 1.4 to 1.5, and the degree of oligomerization was mainly in the range of monomer and dimerization. However, for the products under low temperature condition, the RI value was in the

range of 1.4 to 1.55, and the degree of oligomers could reach tetramerization. As the degree of oligomerization increased, RI was gradually increasing as well. This generally explained the higher RI of SOA under lower temperature condition.

To further validate our speculation and identify the relationship between the measured RI values and the chemical composition of *n*-dodecane SOA, we chosed a surrogate system containing 11 PHA oligomerization and 2 cyclization reactions (with different degree of unsaturation and functional groups) to calculate the expected RI values using the quantitative structure–property relationship (QSPR) method based on the molecular formula and structure (Redmond and

- 290 Thompson, 2011). The details of the 13 reactions and related molecular information were shown in Figure S3 and Table 2, while the relationship between the predicted RI values with the degree of unsaturation, and the degree of oligomerization were shown in Figure 6. Strong correlations were observed between predicted RI values and unsaturation, and the degree of oligomerization. With the same unsaturation, the RI values would increase with the increasing degree of oligomerization; under the same degree oligomerization, the RI values would increase with the increasing degree of unsaturation. In addition,
- 295 the RI values of the surrogate system were similar with the identified substances, which further confirmed the effect of oligomers on RI values.

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In addition to the oligomer reasons above, the difference in wall losses and gas-particle partitioning of gas-phase products might partially contribute to the RI enhancement under low temperature condition. We have measured the wall loss rates of particles under both room and low temperatures, and found that the wall loss rate under low temperature condition (0.0025 ~ 0.0028 min⁻¹) is larger than that under room temperature condition (0.0018 ~ 0.0020 min⁻¹). However, this difference in particle wall loss rate can only slightly change the SOA mass concentration, but not the particle chemical composition. Therefore, it is unlikely to change the optical properties. Nevertheless, the difference in vapor wall loss rates of higher-volatility compounds (while for lower-volatility compounds, their dominant fates are condensation so temperature may affect little on their losses), which may lead to their lower proportions in SOA particles. As their RIs are generally lower than lower-volatility compounds (Li et al., 2018), this proportion change can probably enhance SOA RI at low temperature. Therefore, the difference in wall losses and gas-particle particle particing of gas-phase products might partially contribute to the

3.5 Atmospheric and Climate Implications

RI enhancement under low temperature condition.

- 310 Figure 7 showed the ratios of light extinction efficiency (Q_{ext}) and direct radiative forcing (DRF) of *n*-dodecane SOA under different temperature conditions, from which we could know the impacts of temperature on the role of *n*-dodecane SOA in visibility and radiative balance. As shown in Figure 7a and c, the extinction efficiency (Q_{ext}) of SOA generated under low temperature is larger than SOA generated under room temperature in the size range of 50 to 200 nm, the range of which are the most atmospherically relevant (Zhang et al., 2015; Guo et al., 2014). The enhancement is about 7%-20% at 532 nm, and
- about 1%-21% at 375 nm. This suggests that the extinction efficiency (Q_{ext}) of SOA formed from *n*-dodecane (perhaps other long-chain alkanes in general) was higher in winter than in summer, which would result in lower visibility.

The enhancement in light extinction of SOA and oligomer compositions formed under low temperature condition might provide some possible inspiration for the regional visibility issues, especially the suburban areas. It had been reported that UV-scattering particles in the boundary layer could accelerate photochemical reactions and haze production (Sun et al.,

- 320 2014). The observations above showed that the scattering property of formed SOA increased under low temperature condition, which might provide one possible reason for the rapid occurrence of haze in suburban areas in winter. According to field observations, haze occurred frequently in winter, especially in China (Cheng et al., 2016; Huang et al., 2014; Guo et al., 2014; Parrish et al., 2007). When haze occurred, it was often accompanied by high NO_x, especially in urban areas. The temperature effects under high-NO_x conditions are also important and need to be investigated in future studies.
- As shown in Figure 7b and d, the DRF under low temperature condition was generally larger than room temperature condition: the enhancement was about 6%-19% at 532 nm (50-200 nm), and about 7%-22% at 375 nm (50-180 nm); while for the size range of 180-200 nm, low temperature decreased the DRF ratio about 3% at 375 nm. This phenomenon implied that the SOA generated in winter (low temperature) might have larger DRF on the Earth's surface than the SOA generated in summer (room temperature). It may also imply that the temperature condition should be considered when evaluating the 330 DRF of the aerosol particles generated in the atmosphere.

4 Conclusions

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To the best of our knowledge, this was the first report about optical properties of long-chain alkane SOA at low temperature condition. The modification in temperature significantly changed the chemical composition of the particulate phase. From the oligomer component at low temperature condition, it was presumed that the oligomerization was dominant at low temperature. The presence of oligomers in the SOA particles resulted in an increase of RI values. This study would help to improve our understanding of the lower visibility and the formation of haze in winter. Our results also showed the need for further investigation on the atmospheric parameters influencing SOA formation and optical properties.

Author contributions. WW conceived and led the studies. JL,WZ, and CP performed chamber simulation and data analysis.
 KL, LZ, BS, YC, ML, HL and MG discussed the results and commented on the manuscript. JL prepared the manuscript with contributions from all co-authors.

Data availability. The data used in this study are available upon request from the corresponding author.

345 **Competing interests.** The authors declare that they have no conflict of interest.

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References

Aimanant, S., and Ziemann, P. J.: Chemical mechanisms of aging of aerosol formed from the reaction of n-pentadecane with OH radicals in the presence of NOx, Aerosol Sci. Technol., 47, 979-990, 10.1080/02786826.2013.804621, 2013.
 Atkinson, R., and Arey, J.: Atmospheric degradation of volatile organic compounds, Chem. Rev., 103, 4605-4638,

10.1021/cr0206420, 2003.

- Bohren, C. F. H., D. R.: Absorption and Scattering of Light by Small Particles, 1983.
- 355 Bond, T. C., and Bergstrom, R. W.: Light absorption by carbonaceous particles: An investigative review, Aerosol Sci. Technol., 40, 27-67, 10.1080/02786820500421521, 2006.
 - Boyd, C. M., Nah, T., Xu, L., Berkemeier, T., and Ng, N. L.: Secondary organic aerosol (SOA) from nitrate radical oxidation of monoterpenes: Effects of temperature, dilution, and humidity on aerosol formation, mixing, and evaporation, Environmental Sci. Technol., 51, 7831-7841, 10.1021/acs.est.7b01460, 2017.
- 360 Cappa, C. D., Che, D. L., Kessler, S. H., Kroll, J. H., and Wilson, K. R.: Variations in organic aerosol optical and hygroscopic properties upon heterogeneous OH oxidation, J. Geophys. Res., 116, 10.1029/2011jd015918, 2011.
 - Charron, A., Polo-Rehn, L., Besombes, J.-L., Golly, B., Buisson, C., Chanut, H., Marchand, N., Guillaud, G., and Jaffrezo, J.-L.: Identification and quantification of particulate tracers of exhaust and non-exhaust vehicle emissions, Atmos. Chem. Phys., 19, 5187-5207, 10.5194/acp-19-5187-2019, 2019.
- 365 Cheng, Y., Zheng, G., Wei, C., Mu, Q., Zheng, B., Wang, Z., Gao, M., Zhang, Q., He, K., Carmichael, G., Poschl, U., and Su, H.: Reactive nitrogen chemistry in aerosol water as a source of sulfate during haze events in China, Sci. Adv., 2, e1601530-e1601530, 10.1126/sciadv.1601530, 2016.
- Clark, C. H., Kacarab, M., Nakao, S., Asa-Awuku, A., Sato, K., and Cocker, D. R., 3rd: Temperature effects on secondary organic aerosol (SOA) from the dark ozonolysis and photo-oxidation of isoprene, Environ. Sci. Technol., 10.1021/acs.est.5b05524, 2016.
 - Cui, L., Zhang, Z., Huang, Y., Lee, S. C., Blake, D. R., Ho, K. F., Wang, B., Gao, Y., Wang, X. M., and Louie, P. K. K.: Measuring OVOCs and VOCs by PTR-MS in an urban roadside microenvironment of Hong Kong: relative humidity and temperature dependence, and field intercomparisons, Atmospheric Meas. Tech., 9, 5763-5779, 10.5194/amt-9-5763-2016, 2016.
- 375 Denjean, C., Formenti, P., Picquet-Varrault, B., Katrib, Y., Pangui, E., Zapf, P., and Doussin, J. F.: A new experimental approach to study the hygroscopic and optical properties of aerosols: application to ammonium sulfate particles, Atmos. Meas. Tech., 7, 183-197, 10.5194/amt-7-183-2014, 2014.
 - Denjean, C., Formenti, P., Picquet-Varrault, B., Pangui, E., Zapf, P., Katrib, Y., Giorio, C., Tapparo, A., Monod, A., Temime-Roussel, B., Decorse, P., Mangeney, C., and Doussin, J. F.: Relating hygroscopicity and optical properties to
- 380 chemical composition and structure of secondary organic aerosol particles generated from the ozonolysis of α-pinene, Atmos. Chem. Phys., 15, 3339-3358, 10.5194/acp-15-3339-2015, 2015.
 - Ding, X., Zhang, Y. Q., He, Q. F., Yu, Q. Q., Wang, J. Q., Shen, R. Q., Song, W., Wang, Y. S., and Wang, X. M.: Significant increase of aromatics-derived secondary organic aerosol during fall to winter in China, Environ. Sci. Technol., 51, 7432-7441, 10.1021/acs.est.6b06408, 2017.
- 385 Emanuelsson, E. U., Watne, A. K., Lutz, A., Ljungstrom, E., and Hallquist, M.: Influence of humidity, temperature, and radicals on the formation and thermal properties of secondary organic aerosol (SOA) from ozonolysis of beta-pinene, J. Phys. Chem. A, 117, 10346-10358, 10.1021/jp4010218, 2013.
- Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in cloud droplets and aqueous particles (aqSOA): a review of laboratory, field and model studies, Atmos. Chem. Phys., 11, 11069-11102, 10.5194/acp-11-11069-2011, 2011.
 - 12

- Fahnestock, K. A. S., Yee, L. D., Loza, C. L., Coggon, M. M., Schwantes, R., Zhang, X., Dalleska, N. F., and Seinfeld, J. H.: Secondary organic aerosol composition from C-12 Alkanes, J. Phys. Chem. A, 119, 4281-4297, 10.1021/jp501779w, 2015.
- Flores, J. M., Washenfelder, R. A., Adler, G., Lee, H. J., Segev, L., Laskin, J., Laskin, A., Nizkorodov, S. A., Brown, S. S.,
 and Rudich, Y.: Complex refractive indices in the near-ultraviolet spectral region of biogenic secondary organic aerosol aged with ammonia, Phys. Chem. Chem. Phys., 16, 10629-10642, 10.1039/c4cp01009d, 2014.
- Gentner, D. R., Isaacman, G., Worton, D. R., Chan, A. W. H., Dallmann, T. R., Davis, L., Liu, S., Day, D. A., Russell, L. M., Wilson, K. R., Weber, R., Guha, A., Harley, R. A., and Goldstein, A. H.: Elucidating secondary organic aerosol from diesel and gasoline vehicles through detailed characterization of organic carbon emissions, P. Natl Acad. Sci. USA, 109, 18318-18323, 10.1073/pnas.1212272109, 2012.
- Gentner, D. R., Jathar, S. H., Gordon, T. D., Bahreini, R., Day, D. A., El Haddad, I., Hayes, P. L., Pieber, S. M., Platt, S. M., de Gouw, J., Goldstein, A. H., Harley, R. A., Jimenez, J. L., Prévâ, A. S. H., and Robinson, A. L.: Review of urban secondary organic aerosol formation from gasoline and diesel motor vehicle emissions, Environ. Sci. Technol., 51, 1074-1093, 10.1021/acs.est.6b04509, 2017.
- 405 George, C., Ammann, M., D'Anna, B., Donaldson, D. J., and Nizkorodov, S. A.: Heterogeneous photochemistry in the atmosphere, Chem. Revi., 115, 4218-4258, 10.1021/cr500648z, 2015.
 - Girolami, G. S.: A simple" back of the envelope" method for estimating the densities and molecular volumes of liquids and solids, J. Chem. Educ., 71, 1994.
- Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M., Zeng, L., Molina, M. J., and Zhang,
 R.: Elucidating severe urban haze formation in China, P. Natl Acad. Sci. USA, 111, 17373-17378, 10.1073/pnas.1419604111, 2014.
 - Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A., Prevot, A. S. H., Seinfeld, J. H., Surratt,
- 415 J. D., Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic aerosol: current and emerging issues, Atmos. Chem. Phys., 9, 5155-5236, 2009.
 - Huang, M., Hao, L., Cai, S., Gu, X., Zhang, W., Hu, C., Wang, Z., Fang, L., and Zhang, W.: Effects of inorganic seed aerosols on the particulate products of aged 1,3,5-trimethylbenzene secondary organic aerosol, Atmos. Environ., 152, 490-502, 10.1016/j.atmosenv.2017.01.010, 2017.
- 420 Huang, R. J., Zhang, Y., Bozzetti, C., Ho, K. F., Cao, J. J., Han, Y., Daellenbach, K. R., Slowik, J. G., Platt, S. M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa, M., Ciarelli, G., Piazzalunga, A., Schwikowski, M., Abbaszade, G., Schnelle-Kreis, J., Zimmermann, R., An, Z., Szidat, S., Baltensperger, U., El Haddad, I., and Prevot, A. S.: High secondary aerosol contribution to particulate pollution during haze events in China, Nature, 514, 218-222, 10.1038/nature13774, 2014.
- 425 Huang, W., Saathoff, H., Pajunoja, A., Shen, X., Naumann, K.-H., Wagner, R., Virtanen, A., Leisner, T., and Mohr, C.: αpinene secondary organic aerosol at low temperature: Chemical composition and implications for particle viscosity, Atmos. Chem. Phys., 18, 2883-2898, 2018.
 - Hunter, J. F., Carrasquillo, A. J., Daumit, K. E., and Kroll, J. H.: Secondary organic aerosol formation from acyclic, monocyclic, and polycyclic alkanes, Environ. Sci. Technol., 48, 10227-10234, 10.1021/es502674s, 2014.
- 430 Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., Van Dingenen, R., Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski, Y., Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter, R., Myhre, C. E. L., Tsigaridis, K., Vignati, E., Stephanou, E. G., and Wilson, J.: Organic aerosol and global climate modelling: a review, Atmos. Chem. Phys., 5, 1053-1123, 2005.
- Kim, H., and Paulson, S. E.: Real refractive indices and volatility of secondary organic aerosol generated from photooxidation and ozonolysis of limonene, α-pinene and toluene, Atmos. Chem. Phys., 13, 7711-7723, 10.5194/acp-13-7711-2013, 2013.
 - Kim, H., Liu, S., Russell, L. M., and Paulson, S. E.: Dependence of real refractive indices on O:C, H:C and mass fragments of secondary organic aerosol generated from ozonolysis and photooxidation of limonene and alpha-pinene, Aerosol Sci. Technol., 48, 498-507, 10.1080/02786826.2014.893278, 2014.
- 440 Koss, A. R., Warneke, C., Yuan, B., Coggon, M. M., Veres, P. R., and de Gouw, J. A.: Evaluation of NO⁺ reagent ion

chemistry for online measurements of atmospheric volatile organic compounds, Atmos. Meas. Tech., 9, 2909-2925, 10.5194/amt-9-2909-2016, 2016.

- Lambe, A. T., Onasch, T. B., Croasdale, D. R., Wright, J. P., Martin, A. T., Franklin, J. P., Massoli, P., Kroll, J. H., Canagaratna, M. R., Brune, W. H., Worsnop, D. R., and Davidovits, P.: Transitions from functionalization to fragmentation reactions of laboratory secondary organic aerosol (SOA) generated from the OH oxidation of alkane precursors, Environ, Sci. Technol., 46, 5430-5437, 10.1021/es300274t, 2012.
- Lambe, A. T., Cappa, C. D., Massoli, P., Onasch, T. B., Forestieri, S. D., Martin, A. T., Cummings, M. J., Croasdale, D. R., Brune, W. H., Worsnop, D. R., and Davidovits, P.: Relationship between oxidation level and optical properties of secondary organic aerosol, Environ. Sci. Technol., 47, 6349-6357, 10.1021/es401043j, 2013.
- 450 Lambe, A. T., Chhabra, P. S., Onasch, T. B., Brune, W. H., Hunter, J. F., Kroll, J. H., Cummings, M. J., Brogan, J. F., Parmar, Y., Worsnop, D. R., Kolb, C. E., and Davidovits, P.: Effect of oxidant concentration, exposure time, and seed particles on secondary organic aerosol chemical composition and yield, Atmos. Chem. Phys., 15, 3063-3075, 10.5194/acp-15-3063-2015, 2015.
- Lamkaddam, H., Gratien, A., Pangui, E., Cazaunau, M., Picquet-Varrault, B., and Doussin, J.-F.: High-NOx photooxidation of n-dodecane: Temperature dependence of SOA formation, Environ. Sci. Technol., 10.1021/acs.est.6b03821, 2016.
 - Laskin, A., Laskin, J., and Nizkorodov, S. A.: Chemistry of atmospheric brown carbon, Chem. Rev., 115, 4335-4382, 10.1021/cr5006167, 2015.
 - Lee, A. K. Y., Zhao, R., Li, R., Liggio, J., Li, S.-M., and Abbatt, J. P. D.: Formation of light absorbing organo-nitrogen species from evaporation of droplets containing glyoxal and ammonium sulfate, Environ. Sci. Technol., 47, 12819-12826, 10.1021/es402687w, 2013.
 - Li, J., Li, K., Wang, W., Wang, J., Peng, C., and Ge, M.: Optical properties of secondary organic aerosols derived from longchain alkanes under various NO_x and seed conditions, Sci. Total Environ., 579, 1699-1705, 10.1016/j.scitotenv.2016.11.189, 2017a.
- Li, J., Wang, W., Li, K., Zhang, W., Ge, M., and Peng, C.: Development and application of the multi-wavelength cavity ring-down aerosol extinction spectrometer, J. Environ. Sci. (China), 76, 227-237, 10.1016/j.jes.2018.04.030, 2019.
 - Li, K., Wang, W., Ge, M., Li, J., and Wang, D.: Optical properties of secondary organic aerosols generated by photooxidation of aromatic hydrocarbons, Sci. Rep., 4, 4922-4922, 10.1038/srep04922, 2014.
 - Li, K., Li, J., Liggio, J., Wang, W., Ge, M., Liu, Q., Guo, Y., Tong, S., Li, J., Peng, C., Jing, B., Wang, D., and Fu, P.: Enhanced light scattering of secondary organic aerosols by multiphase reactions, Environ. Sci. Technol., 51, 1285-1292, 10.1021/acs.est.6b03229, 2017b.
 - Li, K., Li, J., Wang, W., Li, J., Peng, C., Wang, D., and Ge, M.: Effects of gas-particle partitioning on refractive index and chemical composition of m-xylene secondary organic aerosol, J. Phys. Chem. A, 10.1021/acs.jpca.7b12792, 2018.
 - Lim, Y. B., and Ziemann, P. J.: Products and mechanism of secondary organic aerosol formation from reactions of n-alkanes with OH radicals in the presence of NO_x, Environ. Sci. Technol., 39, 9229-9236, 10.1021/es051447g, 2005.
- 475 Lim, Y. B., and Ziemann, P. J.: Effects of molecular structure on aerosol yields from OH radical-initiated reactions of linear, branched, and cyclic alkanes in the presence of NO_x, Environ. Sci. Technol., 43, 2328-2334, 10.1021/es803389s, 2009a.
 - Lim, Y. B., and Ziemann, P. J.: Chemistry of secondary organic aerosol formation from OH radical-initiated reactions of linear, branched, and cyclic alkanes in the presence of NO_x, Aerosol Sci. Technol., 43, 604-619, 10.1080/02786820902802567, 2009b.
- 480 Lin, P., Liu, J., Shilling, J. E., Kathmann, S. M., Laskin, J., and Laskin, A.: Molecular characterization of brown carbon (BrC) chromophores in secondary organic aerosol generated from photo-oxidation of toluene, Phys. Chem. Chem. Phys., 17, 23312-23325, 10.1039/c5cp02563j, 2015.
 - Liu, H., Man, H., Cui, H., Wang, Y., Deng, F., Wang, Y., Yang, X., Xiao, Q., Zhang, Q., Ding, Y., and He, K.: An updated emission inventory of vehicular VOCs and IVOCs in China, Atmos. Chem. Phys., 17, 12709-12724, 10.5194/acp-17-12709-2017, 2017.
 - Liu, P. F., Abdelmalki, N., Hung, H. M., Wang, Y., Brune, W. H., and Martin, S. T.: Ultraviolet and visible complex refractive indices of secondary organic material produced by photooxidation of the aromatic compounds toluene and m-xylene, Atmos. Chem. Phys., 15, 1435-1446, 10.5194/acp-15-1435-2015, 2015.
- Loza, C. L., Craven, J. S., Yee, L. D., Coggon, M. M., Schwantes, R. H., Shiraiwa, M., Zhang, X., Schilling, K. A., Ng, N.
 L., Canagaratna, M. R., Ziemann, P. J., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol yields of 12-

445

460

470

485

14

carbon alkanes, Atmos. Chem. Phys., 14, 1423-1439, 10.5194/acp-14-1423-2014, 2014.

- Lu, J. W., Flores, J. M., Lavi, A., Abo-Riziq, A., and Rudich, Y.: Changes in the optical properties of benzo a pyrene-coated aerosols upon heterogeneous reactions with NO₂ and NO₃, Phys. Chem. Chem. Phys., 13, 6484-6492, 10.1039/c0cp02114h, 2011.
- 495 Lu, K., Guo, S., Tan, Z., Wang, H., Shang, D., Liu, Y., Li, X., Wu, Z., Hu, M., and Zhang, Y.: Exploring atmospheric freeradical chemistry in China: the self-cleansing capacity and the formation of secondary air pollution, Natl Sci. Rev., 10.1093/nsr/nwy073, 2018.

McNeill, V. F.: Aqueous organic chemistry in the atmosphere: sources and chemical processing of organic aerosols, Environ. Sci. Technol., 49, 1237-1244, 10.1021/es5043707, 2015.

- 500 Mellouki, A., Wallington, T. J., and Chen, J.: Atmospheric chemistry of oxygenated volatile organic compounds: impacts on air quality and climate, Chem. Rev., 115, 3984-4014, 10.1021/cr500549n, 2015.
 - Michel Flores, J., Bar-Or, R. Z., Bluvshtein, N., Abo-Riziq, A., Kostinski, A., Borrmann, S., Koren, I., Koren, I., and Rudich, Y.: Absorbing aerosols at high relative humidity: linking hygroscopic growth to optical properties, Atmos. Chem. Phys., 12, 5511-5521, 10.5194/acp-12-5511-2012, 2012.
- 505 Moise, T., Flores, J. M., and Rudich, Y.: Optical properties of secondary organic aerosols and their changes by chemical processes, Chem. Rev., 115, 4400-4439, 10.1021/cr5005259, 2015.
 - Nakayama, T., Matsumi, Y., Sato, K., Imamura, T., Yamazaki, A., and Uchiyama, A.: Laboratory studies on optical properties of secondary organic aerosols generated during the photooxidation of toluene and the ozonolysis ofα-pinene, J. Geophys. Res., 115, 10.1029/2010jd014387, 2010.
- 510 Nakayama, T., Sato, K., Matsumi, Y., Imamura, T., Yamazaki, A., and Uchiyama, A.: Wavelength and NO_x dependent complex refractive index of SOAs generated from the photooxidation of toluene, Atmos. Chem. Phys., 13, 531-545, 10.5194/acp-13-531-2013, 2013.
- Parrish, D. D., Stohl, A., Forster, C., Atlas, E. L., Blake, D. R., Goldan, P. D., Kuster, W. C., and de Gouw, J. A.: Effects of mixing on evolution of hydrocarbon ratios in the troposphere, J. Geophys. Res.-Atmos., 112, 10.1029/2006jd007583, 2007.
- Paulsen, D., Dommen, J., Kalberer, M., Preavot, A. S. H., Richter, R., Sax, M., Steinbacher, M., Weingartner, E., and Baltensperger, U.: Secondary organic aerosol formation by irradiation of 1,3,5-trimethylbenzene-NO_x-H₂O in a new reaction chamber for atmospheric chemistry and physics, Environ. Sci. Technol., 39, 2668-2678, 2005.
- Peng, C., Wang, W., Li, K., Li, J., Zhou, L., Wang, L., and Ge, M.: The optical properties of limonene secondary organic
 aerosols: the role of NO₃, OH and O₃ in the oxidation processes, J. Geophys. Res.: Atmos., 10.1002/2017jd028090, 2018.
 - Phillips, S. M., and Smith, G. D.: Light absorption by charge transfer complexes in brown carbon aerosols, Environ. Sci. Technol. Let., 1, 382-386, 10.1021/ez500263j, 2014.
- Poschl, U.: Atmospheric aerosols: Composition, transformation, climate and health effects, Angew. Chem.-Int. Edit., 44, 7520-7540, 10.1002/anie.200501122, 2005.
- Poschl, U., and Shiraiwa, M.: Multiphase chemistry at the atmosphere-biosphere interface influencing climate and public health in the anthropocene, Chem. Rev., 115, 4440-4475, 10.1021/cr500487s, 2015.

530

- Presto, A. A., Miracolo, M. A., Kroll, J. H., Worsnop, D. R., Robinson, A. L., and Donahue, N. M.: Intermediate-volatility organic compounds: A potential source of ambient oxidized organic aerosol, Environ. Sci. Technol., 43, 4744-4749, 10.1021/es803219q, 2009.
- Price, D. J., Kacarab, M., Cocker, D. R., Purvis-Roberts, K. L., and Silva, P. J.: Effects of temperature on the formation of secondary organic aerosol from amine precursors, Aerosol Sci. Technol., 50, 1216-1226, 10.1080/02786826.2016.1236182, 2016.
- Qi, L., Nakao, S., Tang, P., and R., C. I. D.: Temperature effect on physical and chemical properties of secondary organic aerosol from m-xylene photooxidation, Atmos. Chem. Phys., 10, 3847-3854, 2010.
 - Qing Mu, M. S., Mega Octaviani, Nan Ma, Aijun Ding, Hang Su,, and Gerhard Lammel, U. P., Yafang Cheng: Temperature effect on phase state and reactivity controls atmospheric multiphase chemistry and transport of PAHs, Sci. Adv., 4, 10.1126/sciadv.aap7314, 2018.

Redmond, H., and Thompson, J. E.: Evaluation of a quantitative structure-property relationship (QSPR) for predicting mid-

540 visible refractive index of secondary organic aerosol (SOA), Phys. Chem. Chem. Phys. : PCCP, 13, 6872-6882,

10.1039/c0cp02270e, 2011.

580

590

- Roy, A., and Choi, Y.: Effect of ambient temperature on species lumping for total organic gases in gasoline exhaust emissions, Atmos. Environ., 152, 240-245, 10.1016/j.atmosenv.2016.11.057, 2017.
- Sareen, N., Waxman, E. M., Turpin, B. J., Volkamer, R., and Carlton, A. G.: Potential of aerosol liquid water to facilitate
 organic aerosol formation: Assessing knowledge gaps about precursors and partitioning, Environ. Sci. Technol., 51, 3327-3335, 10.1021/acs.est.6b04540, 2017.
 - Shi, B., Wang, W., Zhou, L., Sun, Z., Fan, C., Chen, Y., Zhang, W., Qiao, Y., Qiao, Y., and Ge, M.: Atmospheric oxidation of C10~14 n-alkanes initiated by Cl atoms: Kinetics and mechanism, Atmos. Environ., 117166, https://doi.org/10.1016/j.atmosenv.2019.117166, 2019.
- 550 Shrivastava, M., Cappa, C. D., Fan, J., Goldstein, A. H., Guenther, A. B., Jimenez, J. L., Kuang, C., Laskin, A., Martin, S. T., Ng, N. L., Petaja, T., Pierce, J. R., Rasch, P. J., Roldin, P., Seinfeld, J. H., Shilling, J., Smith, J. N., Thornton, J. A., Volkamer, R., Wang, J., Worsnop, D. R., Zaveri, R. A., Zelenyuk, A., and Zhang, Q.: Recent advances in understanding secondary organic aerosol: Implications for global climate forcing, Rev. Geophys., 55, 509-559, 10.1002/2016rg000540, 2017.
- 555 Song, C., Gyawali, M., Zaveri, R. A., Shilling, J. E., and Arnott, W. P.: Light absorption by secondary organic aerosol from alpha-pinene: Effects of oxidants, seed aerosol acidity, and relative humidity, J. Geophys. Res.-Atmos., 118, NIL_306-NIL_314, 2013.
- Sun, Y., Jiang, Q., Wang, Z., Fu, P., Li, J., Yang, T., and Yin, Y.: Investigation of the sources and evolution processes of severe haze pollution in Beijing in January 2013, J. Geophys. Res.-Atmos., 119, 4380-4398, 10.1002/2014jd021641, 2014.
- Svendby, T. M., Lazaridis, M., and Tørseth, K.: Temperature dependent secondary organic aerosol formation from terpenes and aromatics, J. Atmos. Chem., 59, 25-46, 10.1007/s10874-007-9093-7, 2008.
 - Takekawa, H., Minoura, H., and Yamazaki, S.: Temperature dependence of secondary organic aerosol formation by photooxidation of hydrocarbons, Atmos. Environ., 37, 3413-3424, 10.1016/s1352-2310(03)00359-5, 2003.
- 565 Titos, G., Cazorla, A., Zieger, P., Andrews, E., Lyamani, H., Granados-Muñoz, M. J., Olmo, F. J., and Alados-Arboledas, L.: Effect of hygroscopic growth on the aerosol light-scattering coefficient: A review of measurements, techniques and error sources, Atmos. Environ., 141, 494-507, 10.1016/j.atmosenv.2016.07.021, 2016.
- Trainic, M., Abo Riziq, A., Lavi, A., Flores, J. M., and Rudich, Y.: The optical, physical and chemical properties of the products of glyoxal uptake on ammonium sulfate seed aerosols, Atmos. Chem. Phys., 11, 9697-9707, 10.5194/acp-11-9697-2011, 2011.
- von Schneidemesser, E., Monks, P. S., Allan, J. D., Bruhwiler, L., Forster, P., Fowler, D., Lauer, A., Morgan, W. T., Paasonen, P., Righi, M., Sindelarova, K., and Sutton, M. A.: Chemistry and the linkages between air quality and climate change, Chem. Rev., 115, 3856-3897, 10.1021/acs.chemrev.5b00089, 2015.

Wang, J. M., Jeong, C.-H., Zimmerman, N., Healy, R. M., Hilker, N., and Evans, G. J.: Real-world emission of particles
 from vehicles: Volatility and the effects of ambient temperature, Environ. Sci. Technol., 10.1021/acs.est.6b05328, 2017.

Wang, L., Wang, W., and Ge, M.: Extinction efficiencies of mixed aerosols measured by aerosol cavity ring down spectrometry, Chinese Sci. Bull., 57, 2567-2573, 10.1007/s11434-012-5146-7, 2012.

Wang, W.-G., Li, K., Zhou, L., Ge, M.-F., Hou, S.-Q., Tong, S.-R., Mu, Y.-J., and Jia, L.: Evaluation and application of dual-reactor chamber for studying atmospheric oxidation processes and mechanisms, Acta Phys.-Chim. Sin., 31, 1251-1259, 10.3866/pku.whxb201504161, 2015.

Warren, B., Austin, R. L., and Cocker, D. R.: Temperature dependence of secondary organic aerosol, Atmos. Environ., 43, 3548-3555, 10.1016/j.atmosenv.2009.04.011, 2009.

Ye, Q., Robinson, E. S., Ding, X., Ye, P., Sullivan, R. C., and Donahue, N. M.: Mixing of secondary organic aerosols versus relative humidity, P. Natl. Acad. Sci. USA, 113, 12649-12654, 10.1073/pnas.1604536113, 2016.

585 Yee, L. D., Craven, J. S., Loza, C. L., Schilling, K. A., Ng, N. L., Canagaratna, M. R., Ziemann, P. J., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol formation from low-NO(x) photooxidation of dodecane: evolution of multigeneration gas-phase chemistry and aerosol composition, J. Phys. Chem. A, 116, 6211-6230, 10.1021/jp211531h, 2012.

Yee, L. D., Craven, J. S., Loza, C. L., Schilling, K. A., Ng, N. L., Canagaratna, M. R., Ziemann, P. J., Flagan, R. C., and Seinfeld, J. H.: Effect of chemical structure on secondary organic aerosol formation from C12 alkanes, Atmos. Chem.

Phys., 13, 11121-11140, 10.5194/acp-13-11121-2013, 2013.

- Zhang, R., Wang, G., Guo, S., Zamora, M. L., Ying, Q., Lin, Y., Wang, W., Hu, M., and Wang, Y.: Formation of urban fine particulate matter, Chem. Rev., 115, 3803-3855, 10.1021/acs.chemrev.5b00067, 2015.
- Zhao, B., Wang, S., Donahue, N. M., Jathar, S. H., Huang, X., Wu, W., Hao, J., and Robinson, A. L.: Quantifying the effect of organic aerosol aging and intermediate-volatility emissions on regional-scale aerosol pollution in China, Sci. Rep., 6, 28815, 10.1038/srep28815, 2016.
- Zhao, Y., Hennigan, C. J., May, A. A., Tkacik, D. S., de Gouw, J. A., Gilman, J. B., Kuster, W. C., Borbon, A., and Robinson, A. L.: Intermediate-volatility organic compounds: a large source of secondary organic aerosol, Environ. Sci. Technol., 48, 13743-13750, 10.1021/es5035188, 2014.
- 600 Zhao, Z., Le, C., Xu, Q., Peng, W., Jiang, H., Lin, Y.-H., Cocker, D. R., and Zhang, H.: Compositional evolution of secondary organic aerosol as temperature and relative humidity cycle in atmospherically relevant ranges, ACS Earth Space Chem., 3, 2549-2558, 10.1021/acsearthspacechem.9b00232, 2019.
 - Zhong, M., and Jang, M.: Dynamic light absorption of biomass-burning organic carbon photochemically aged under natural sunlight, Atmos. Chem. Phys., 14, 1517-1525, 10.5194/acp-14-1517-2014, 2014.
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Experiments No. ^a	HC (ppb)	H ₂ O ₂ (ppm)	NO _X (ppb)	RH (%)	hamber experiments. Temperature (°C)	<mark>Mass^b (µg/m³)</mark>
Dod-R-1	58	1.03	<1	<5	25	<mark>155</mark>
Dod-R-2	52	1.03	<1	<5	25	<mark>135</mark>
Dod-L-1	43	1.07	<1	<5	5	<mark>128</mark>
Dod-L-2	50	1.09	<1	<5	5	<mark>133</mark>

^aExperimental conditions: *n*-Dodecane room temperature (Dod-R), *n*-Dodecane low temperature (Dod-L)

^bThe mass concentration are obtained without wall loss correction.

)	Table 2: The calculated RI values for the surrogate system: oligomers (11 types of PHA) and furan derivatives.							
	Molecular Formula	Molecular Weight	Polarizability	Unsaturation	Predicted RI			
	$C_{12}H_{26}O_3$	218.32	24.57	0	1.39017			
	C ₁₂ H ₂₄ O ₃	216.31	24.23	1	1.42236			
	C12H22O4	230.29	24.46	2	1.45422			
	C12H24O4	232.31	24.8	1	1.42203			
	C12H22O5	246.29	25.03	2	1.45389			
	C12H24O5	248.3	25.37	1	1.42174			
	C12H20O5	244.27	24.69	3	1.48607			
	$C_{12}H_{18}O_{6}$	258.26	24.92	4	1.51681			
	$C_{12}H_{20}O_{6}$	260.27	25.26	3	1.48574			
	$C_{12}H_{24}O_{6}$	264.3	25.94	1	1.4215			
	C12H22O6	262.29	25.6	2	1.4536			
	C22H44O6	404.56	44.44	1	1.43921			
	C22H46O6	406.58	44.78	0	1.40724			
	C22H42O7	418.54	44.67	2	1.47106			

C ₂₂ H ₄₄ O ₇	420.56	45.01	1	1.43909
C22H42O8	434.54	45.24	2	1.47095
C22H44O8	436.56	45.58	1	1.43899
C22H40O8	432.53	44.9	3	1.50292
C ₂₂ H ₃₈ O ₉	446.51	45.13	4	1.53478
C22H40O9	448.53	45.47	3	1.50281
C22H44O9	452.56	46.15	1	1.43948
C22H42O9	450.54	45.81	2	1.47085
C32H64O9	592.82	64.65	1	1.45319
C32H66O9	594.83	64.99	0	1.42134
C32H62O10	606.8	64.88	2	1.48505
C ₃₂ H ₆₄ O ₁₀	608.82	65.22	1	1.4532
C ₃₂ H ₆₂ O ₁₁	622.8	65.45	2	1.48505
C32H64O11	624.81	65.79	1	1.45361
C ₃₂ H ₆₀ O ₁₁	620.78	65.11	3	1.5169
C ₃₂ H ₅₈ O ₁₂	634.77	65.34	4	1.54876
C ₃₂ H ₆₀ O ₁₂	637.78	65.68	3	1.51691
C32H64O12	640.81	66.36	1	1.45403
C32H62O12	638.8	66.02	2	1.48547
C42H84O12	781.07	84.86	1	1.46637
C42H86O12	783.09	85.2	0	1.4349
C42H82O13	795.05	85.09	2	1.49823
C42H84O13	797.07	85.43	1	1.46676
C42H82O14	811.05	85.66	2	1.49861
C42H84O14	813.07	86	1	1.46715
C42H80O14	809.04	85.32	3	1.53008
C42H78O15	823.02	85.55	4	1.56194
C42H80O15	825.04	85.89	3	1.53047
C42H84O15	829.07	86.57	1	1.46753
C42H82O15	827.05	86.23	2	1.499
C12H23O2	199.3	23.49	1.5	1.43592
C12H21O3	213.28	23.72	2.5	1.46778
C12H23O3	215.3	24.06	1.5	1.43562
C12H22O3	214.29	23.89	2	1.45228
C12H21O4	229.28	24.29	2.5	1.46748
C12H20O4	228.27	24.12	3	1.48413

$C_{12}H_{22}O_4$	230.29	24.46	2	1.45198
C12H19O4	227.27	23.95	3.5	1.49964
C12H17O5	241.25	24.18	4.5	1.5315
C12H19O5	243.27	24.52	3.5	1.49934
$C_{12}H_{18}O_5$	242.26	24.35	4	1.51599
C12H22O5	246.29	25.03	2	1.45173
C12H21O5	245.28	24.86	2.5	1.46723
C12H20O5	244.27	24.69	3	1.48384

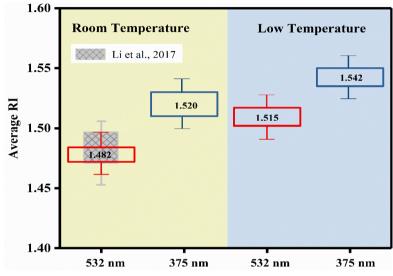


Figure 1: Summary of the averaged RI values of dodecane SOA under room temperature 25 °C (the light orange area) and low temperature 5 °C (the light blue area) in the wavelength of 532 nm and 375 nm. The red box is the averaged RI value for *n*-dodecane in 532 nm, the shaded boxes are the RI values from our previous study in 532 nm (Li et al., 2017a), the blue box is the

615 dodecane in 532 nm, the shaded boxes are th averaged RI value for *n*-dodecane in 375 nm.

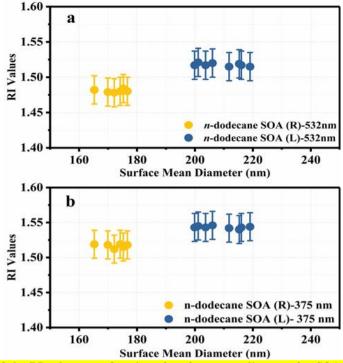


Figure 2: Variation tendency of the RI values as a function of surface mean diameter for SOA produced under room and low

620 temperature conditions at (a) 532 nm and (b) 375 nm.

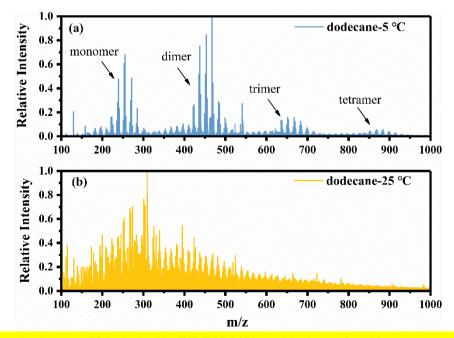
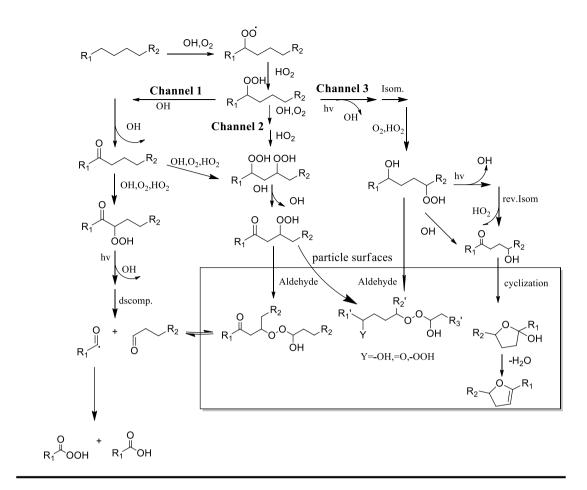


Figure 3: Mass spectra of *n*-dodecane SOA obtained by ESI-TOF-MS in positive ion mode. (a) low temperature condition; (b) room temperature condition.



Low Temperature, Particle Surface

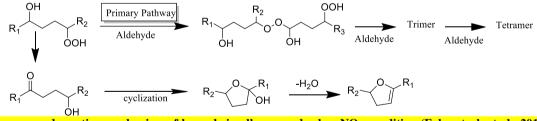
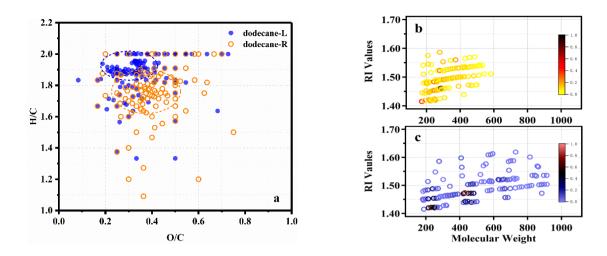
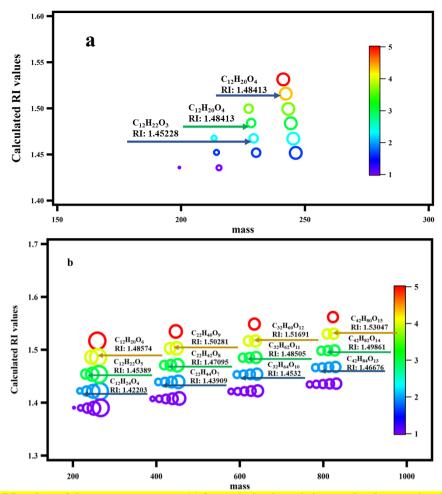




Figure 4: The proposed reaction mechanism of long-chain alkanes under low-NOx condition (Fahnestock et al., 2015;Yee et al., 2013, 2012)



630 Figure 5: (a) Van Krevelen plots showing O/C and H/C ratios for identified products by EST-TOF-MS: the orange circle is for the SOA generated under room temperature condition; the blue circle is for the SOA generated under low temperature condition; (b) calculated RI values of the selected identified molecular species from MS spectra under room temperature, the color map refers to the relative intensity of the molecular formula; (c) calculated RI values of the selected identified molecular species from MS spectra under low temperature, the color map refers to the relative intensity of the molecular species from MS spectra under low temperature.



660 Figure 6: Calculated RIs values of the surrogate system: (a) furan derivatives via intramolecular reaction (b) the oligomers via the formation of peroxy-hemiacetal. The color map refers to the unsaturation of the molecular formula. The size of the circle refers to the O/C ratio.

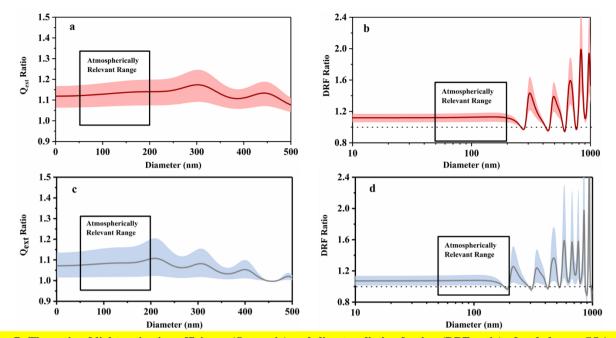


Figure 7: The ratio of light extinction efficiency (Q_{ext} ratio) and direct radiative forcing (DRF ratio) of *n*-dodecane SOA under different temperature conditions. (a) Q_{ext} ratio at 532 nm; (b) DRF ratio at 532 nm; (c) Q_{ext} ratio at 375 nm; (d) DRF ratio at 375 nm. The solid line is the average value of the ratio, and the shaded area is the uncertainty.