

Response to Referee #1

The authors thank the reviewer for the comments that improve the quality of the paper. The detailed responses are given as follows. The reviewer comments are shown in italic font, the responses are in regular font, and the revised text is in bold font.

The authors present model results using the WRF/CMAQ for SOA formation in China. The model is updated for the partitioning of water vapor to the organic aerosol and the non-ideality of the organic phase. A comparison of the model to observations is performed for multiple sites. SOA enhancement during the summer and winter is discussed for the different China domains. The effect of aerosol liquid water on SOA formation and aerosol optical depth are presented. Correlation of the calculated OA hygroscopicity based on the k-Köhler theory to the O:C ratio is performed to show seasonal and multicity variations. Although the publication could provide valuable insights into the factors that govern SOA formation in China it currently lacks a detailed discussion and validation of the presented results. Therefore, the publication requires major revisions suggested below.

Major Comments

Comment 1: *The manuscript is hard to read. Discussion and Results are not well separated in the manuscript. Many phrases are not clear and require more elaboration and better use of English. The table and most figures are poorly made and the science is hard to follow. Examples are given below.*

Response 1: We thank the reviewer for pointing this out. We have updated all the figures and tables by showing monthly-averaged results instead of the daily maximum average as used in the previous version. Also, the text has been revised carefully to make it clear and easy to follow.

Comment 2: *A major drawback in this work is that the model is not capturing the SOA formation during the winter that has been shown to be the dominant organic aerosol source in multiple publications for different domains of China. The authors only in a sentence discuss that the conversion of the POA to SOA may be the reason for these discrepancies but they have no observations to back this up.*

Response 2: We agree with the reviewer that there is no direct observational evidence for this particular episode that POA aging played a significant role in the SOA formation. However, this process is one of the important missing

sources of SOA in several models, field and chamber studies (Robinson et al., 2007; Shrivastava et al., 2008; Zhao et al., 2016; Jimenez et al., 2009; Hodzic et al., 2010; Murphy et al., 2017). Also, organic compounds with intermediate volatility (IVOCs) between SVOCs and traditional VOCs, especially from combustion sources might contribute to SOA as well (Robinson et al., 2007; Shrivastava et al., 2008; Tkacik et al., 2012; Zhao et al., 2014; Zhao et al., 2016; Hodzic et al., 2010), but with high uncertainties in the emission inventory and SOA yields (Shrivastava et al., 2008; Tkacik et al., 2012). We performed an additional simulation using the latest version of the CMAQ model (v5.3.1), which includes a parameterization of these processes. The predicted SOA indeed increases significantly (Figure R1). We further analyzed the contribution of traditional POA (as semi-volatile POA in CMAQv5.3.1), SOA from the oxidation of the semi-volatile POA, traditional SOA, and a new SOA surrogate (pcSOA) representing missing SOA from IVOC oxidation, multigenerational aging of VOC oxidation products, and underestimate of SOA yield due to chamber wall losses (Murphy et al., 2017), finding pcSOA dominates in Beijing and Guangzhou (Figure R2) as well as the whole domain (Figure R1). The averaged SOA/POA ratio in Beijing is increased to 0.83, which is more consistent with field measurements (Zhao et al., 2019; Sun et al., 2016; Sun et al., 2013). However, the emission factors and oxidation rate of pcSOA precursors are highly uncertain and the contribution of pcSOA requires more observational constrains (Murphy et al., 2017).

We also examined the sensitivity of SOA and organic liquid water (ALW_{org}) to pcSOA and POA in an offline calculation in Beijing, Guangzhou, Jinan, and Nanjing. POA has the same properties as we used in the model. Non-volatile isoprene SOA is taken to represent pcSOA as their similarities in saturation vapor pressure and O:C ratio. We found that both SOA and ALW_{org} are positively correlated with pcSOA, increased by 2-5 times in different locations when pcSOA increased by 2 times. The impacts of water partitioning into OPM and non-ideality of the organic-water mixture by including the above process should be explored in a future study.

We extend the discussion of potential reasons for underestimated SOA in the revised text (L309-326):

“Again, no apparent changes of SOA nor OA are observed between case S3 and BS (not shown), since POA is predicted to be the primary contributor to OA at Beijing in winter in the current model, with an averaged SOA/POA ratio of 0.12. This ratio is much lower than the field observation of about 0.45-1.94 (Zhao et al., 2019; Sun et al., 2013; Sun et al., 2016). The bias might be due to the missing SOA converted by partitioning and aging of semi-volatile POA as well as oxidation from

intermediate volatile organic compounds (IVOCs) and VOC oxidation products. Those pathways are shown to be important for SOA formation by modeling, field and chamber studies (Hodzic et al., 2010; Jimenez et al., 2009; Murphy et al., 2017; Robinson et al., 2007; Shrivastava et al., 2008; Tkacik et al., 2012; Zhao et al., 2014; Zhao et al., 2016a).

A sensitivity test was performed by using the newest CMAQ model version 5.3.1 that includes all the above processes in the aerosol module. The SOA/POA ratio in Beijing is improved greatly to be 0.83 in winter. However, high uncertainties still exist in the emissions of the involved precursors and characterization of SOA formation through these processes, needing further constrains by observations. Their influences on water partitioning into OPM and non-ideality of the organic-water mixture on SOA will be evaluated in a future study.”

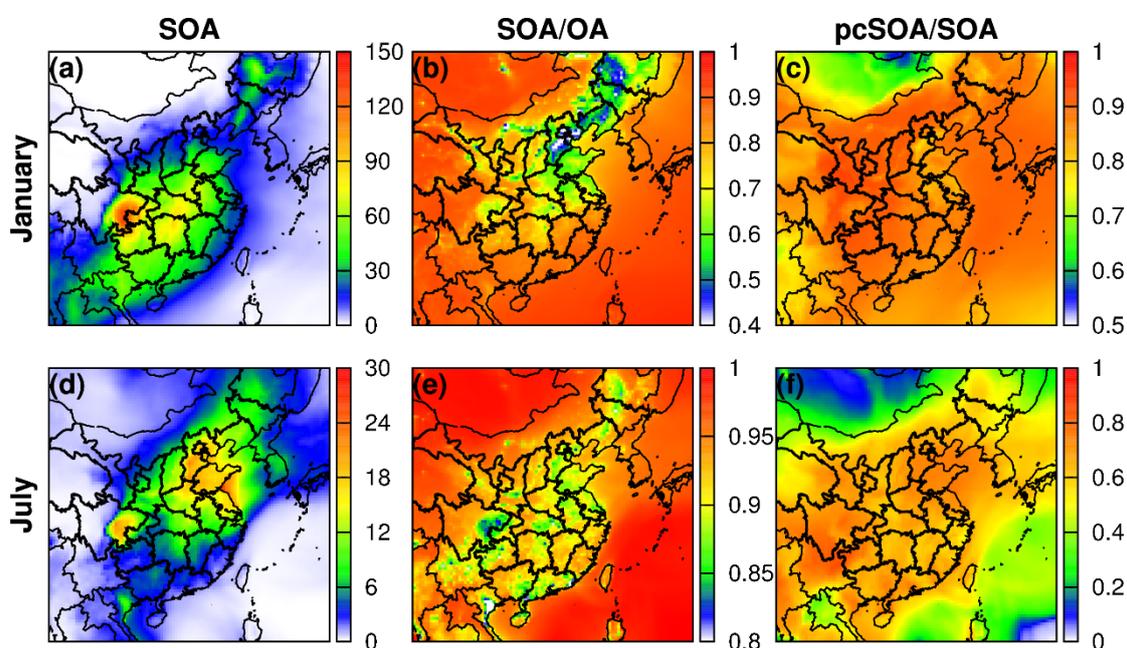


Figure R1. Mean SOA, SOA/OA, and pcSOA/SOA ratio predicted during January and July of 2013 by CMAQv5.3.1.

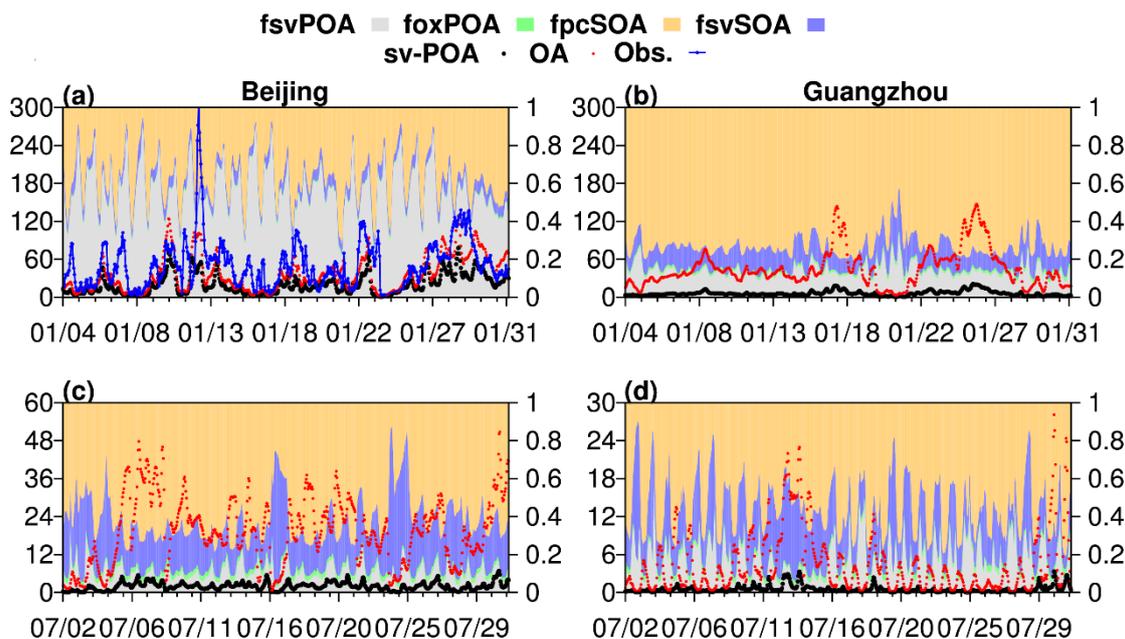


Figure R2. Modeled concentration of semi-volatile POA (sv-POA), OA and fraction of each organic aerosol component fsvPOA (sv-POA), foxPOA (oxidation of sv-POA), fpcSOA (pc-SOA) and fsvSOA (traditional SOA) in Beijing (a, c) and Guangzhou (b, d). Observations of OA in January 2013 at Beijing (Obs.) are also included in (a). The left axis is for the concentrations ($\mu\text{g m}^{-3}$) and the right axis is the fraction of OA components.

Comment 3: *There is no discussion on the influence of nitrate on aerosol liquid water. What fraction of the ALW is related to nitrate and what to organics? How could the ALW related to nitrate influence the partitioning of organics?*

Response 3: In the current model, we separately treated the liquid water associated with organics (ALW_{org}) and inorganics (ALW_{ing}) in the condensed phase. Nitrate is assumed to only affect the inorganic aerosols and ALW_{ing} as the interactions between inorganic and organic phases are not considered currently. This is the same approach used by Pankow et al. (2015). When considering the interactions between inorganic and organic aerosols in a CMAQ model, Pye et al. (2017) found an increase in SOA, which impacts are less significant than the separate treatment of the two phases. However, the interactions among condensed organics, i.e. the polarity of organics in the aerosol were ignored in their study. The interactions between inorganic and organic phases on ALW and SOA are beyond the scope of the current study and will be investigated in the future. We added a statement in the revised text (L207-L208) to make it clear:

“In the current model, we assumed no interactions between the inorganic and organic phases.”

Comment 4: *A comparison of the model to observations should be performed and presented in the main text for both seasons in more detail. The effect of the improvements performed for the SAPRC-11 model is not discussed. The processes added e.g., the heterogeneous formation of nitrate and sulfate on the particle surface, SOA from isoprene, and dicarbonyls surface-controlled reactive uptake are not discussed. What is the effect of these added processes to the overall performance of the model? Each addition and the effect should be discussed in detail in order to support the importance of including them.*

Response 4: The contribution of heterogeneous formation of nitrate and sulfate are not related to SOA formation. Papers documenting these changes have already been cited in the manuscript (Ying et al., 2014; Hu et al., 2016), and the impacts of the heterogeneous chemistry on nitrate and sulfate were discussed in those and another study (Zheng et al., 2015). The predicted nitrate and sulfate have been extensively compared with observations (Shi et al., 2017; Qiao et al., 2018). The improvements in the modeled SOA such as the reactive uptake of dicarbonyls and the isoprene generated epoxydiols have been discussed in previous studies by Ying et al. (2015), Li et al. (2015), Hu et al. (2017) and Liu et al. (2020) and have been shown to greatly increase the predicted SOA concentrations. The focus of this work is partitioning of water into OPM and the polarity of condensed organic compounds on SOA formation in China, which have not been examined so far. We have revised the manuscript to make it clear (L149-L161):

“Model configurations were largely based on that used by Hu et al. (2016) as summarized below. Firstly, SAPRC-11 was expanded for a more detailed treatment of isoprene oxidation and tracking dicarbonyl products (glyoxal and methylglyoxal) from different groups of major precursors (Ying et al., 2015). Secondly, SOA from isoprene epoxydiols (IEPOX), methacrylic acid epoxide (MAE) and dicarbonyls through surface-controlled irreversible reactive uptake were added (Hu et al., 2017; Li et al., 2015; Liu et al., 2020; Ying et al., 2015). Thirdly, the heterogeneous formation of secondary nitrate and sulfate from NO₂ and SO₂ reaction on the particle surfaces (Ying et al., 2014) were added, which is an important source of secondary inorganic aerosols (Zheng et al., 2015) and improves model estimates of nitrate and sulfate (Qiao et al., 2018; Shi et al., 2017). Fourthly, SOA yields were corrected for vapor wall loss (Zhang et al., 2014). Impacts of the above updates on model performances have been extensively discussed in the cited work and will not be further investigated in the current study.”

Comment 5: *MEGAN has been shown to overestimate the isoprene emissions. Would this have a major effect on SOA formation in this work?*

Response 5: The MEGAN model has been shown to overestimate emissions of isoprene in the eastern and southeastern US (Wang et al., 2017; Kota et al., 2015), which is mainly due to databases of emission factors used. While our previous study indeed showed up to 5 times higher isoprene concentrations compared to the observations in Nanjing, the isoprene oxidation products (MACR and MVK) agree well with observations. Another study of Southern China also showed up to 2 times higher of MEGAN compared to measured isoprene fluxes (Situ et al., 2014). Thus, it is still inconclusive whether isoprene was indeed overestimated. If this isoprene overestimation was prevalent throughout the country and the isoprene SOA changed linearly with isoprene emissions, the actual SOA concentration and impacts on SOA due to water partitioning will decrease by 40-50% and 20-30%, respectively. Emissions of isoprene and other biogenic emissions are low during winter so that little or no impact is expected for winter.

Comment 6: *Table 1: There is no discussion of the table in the main text. Abbreviations are not included in the caption. What is MB, GE, Num? Discussion on more statistically relevant values would be beneficial, e.g. what is the R2 of the comparisons? What are the presented values? Medians? Means? What is the domain of the model and how much do the values fluctuate around the domain? What are the uncertainties of the measurements and the model? A figure of the comparison of temperature and relative humidity of obs. vs model with the associated errors and linear regression analysis with the statistics would be informative.*

Response 6: Table 1 shows the mean observation (OBS), mean prediction (PRE), mean bias (MB), gross error (GE), and the number of valid data of temperature and relative humidity in 8 sub-regions of the domain as shown in Figure S1. The table has been revised with explanations of all the abbreviations. A new table (Table S5) has been added to explain each region in Table 1. Since there are too many observation sites in the domain to show the uncertainties and regression information, we added R in the revised Table 1 and updated the table. More information about measurement methodology and uncertainties associated with measurements can be found at the NCDC data website. We discussed Table 1 in the original text (L245-L249):

“Table 1 shows the comparison of WRF predictions and observations in 8 sub-regions of the domain (Figure S1). Observed data are accessible from the National Climatic Data Center at <ftp://ftp.ncdc.noaa.gov/pub/data/noaa>. Temperature and RH are well captured by WRF in YRD, the Pearl River Delta (PRD), and central regions of China (the major regions of eastern China).”

We expanded this discussion in the revised manuscript in L277-L288:

“Temperature and RH are the two meteorological factors that affect SOA formation. Table 1 lists model statistics of mean observation (OBS), mean prediction (PRE), mean bias (MB), gross error (GE) and correlation coefficient (R) based on WRF and observations at monitoring sites located in 8 sub-regions of the domain (Figure S1) during January and July of 2013. The benchmarks for the MM5 model (another meteorology model) of 4-12km horizontal resolution suggested by Emery et al. (2001) are also listed in the table. Details of monitoring sites in the 8 sub-regions are listed in Table S5. Overall, WRF tends to underestimate both temperature and RH. The model shows better agreement with observed temperature as R is higher than that of RH. Both temperature and RH are well captured by the model in YRD, the Pearl River Delta (PRD), and the central regions of China (the major regions of eastern China). In these regions, MB and GE of temperature are -1.2~0.7 K and 1.8~2.6 K, respectively, which are -11.8~5.6% and 9.2~16.8% for RH, respectively.”

Comment 7: *Figure 1: (1) The current Figure has no information regarding the season the measurements are performed. A comparison of both seasons should be made and a figure like Fig1(a) should be made for each season. The time series should include the same site for the two seasons. (2) Why is the base case exactly the same as the S3 in Figure 1(b)? (3) Add errors to the measurements.*

Response 7: (1) Due to limited observations of the simulated episode, we only have OC and OA measurements in January of 2013 at these sites. We used surface PM_{2.5} alternatively to evaluate model performances in July of 2013. As a significant fraction of PM_{2.5} in July is secondary, this still provides an indirect assessment of the model prediction of the oxidation capacity of the atmosphere, which is important for SOA formation. (2) The insignificant difference between BS and S3 in Figure 1(b) is likely due to a much smaller fraction of SOA compared to POA at this location predicted in the current model. Related discussions can refer to Response 2. (3) Unfortunately, there is no error information available for those measurements at this point.

Comment 8: *Figure 2: Why do you use in (a) the base case and not the S3 case? I would consider promoting the updated S3 on the left and the changes on the middle and right panels.*

Response 8: We have replaced case BS with S3 in Figure 2(a) and (d) and showed monthly-averaged results instead of averaged daily maximum in (b)-(c) and (e)-(f) in Figure 2.

Comment 9: *Figure 3: I don't see the point in presenting the ratio of LWC_{org} to SOA. If both are expected to increase during pollution episodes then the ratio might stay the same therefore providing no valuable information. I would plot the SOA to OA as an alternative option or the SOA alone.*

Response 9: One effect of water in the organic phase is that it decreases the average molecular weight of the absorbing organic phase, which could affect the subsequent partitioning of other semi-volatile organic compounds (see Eq 2). The ALW_{org}/SOA allows the readers to see the importance to consider ALW_{org} in the partitioning calculation as it can account for a significant fraction of SOA and lead to a reduced average molecular weight. We agree with the reviewer that this ratio might not be very different between clean and polluted episodes. However, it is nonetheless useful in assessing the importance of including ALW_{org} in SOA modeling. SOA/OA is certainly very useful information as well. Since POA is still significantly higher than SOA, especially during the winter month, SOA/OA ratio did not change significantly when ALW_{org} is considered as shown in Figure R3. We have added this figure in the revised supplemental materials as Figure S13.

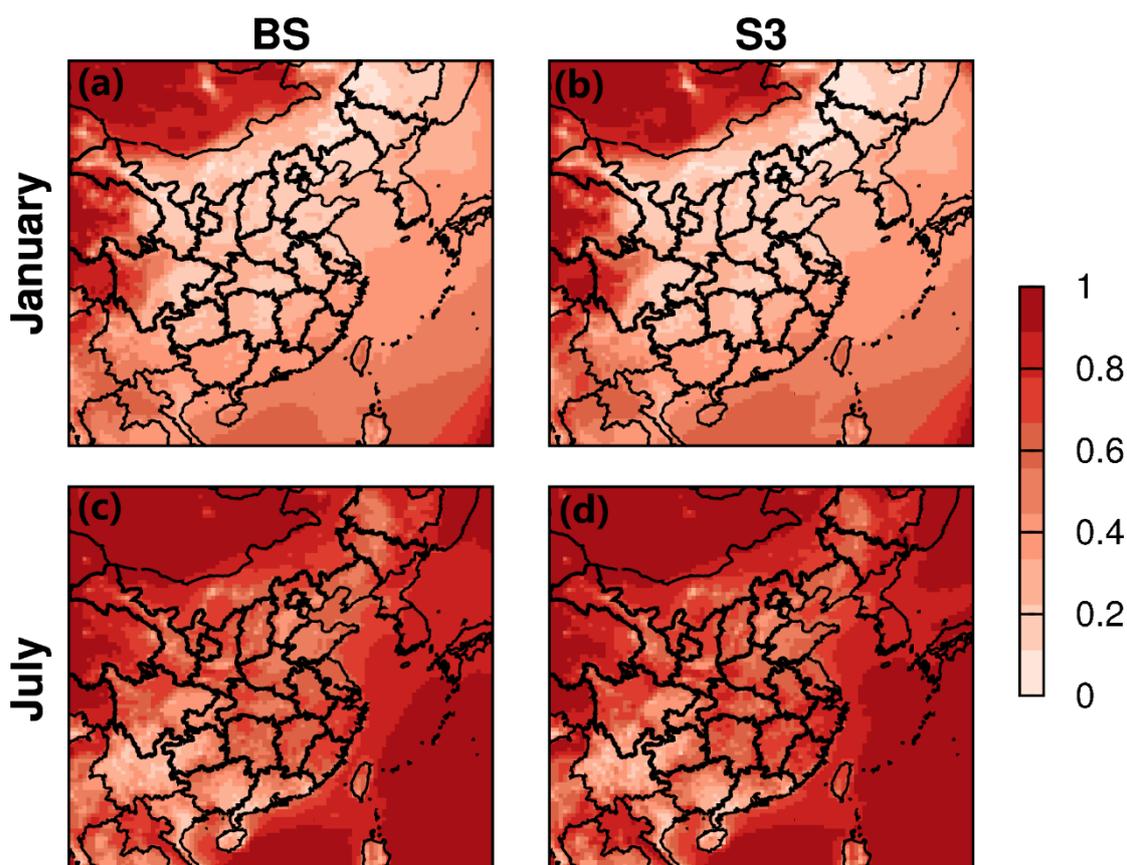


Figure R3. Averaged SOA/OA ratio from case BS and S3 during January and July of 2013.

Comment 10: *Figure 4: The data are really hard to observe. Please change*

colors and increase font size.

Response 10: Figure 4 (as shown below) has been revised to show all the data from each city in each month. Detailed results of each city are shown in Figure S9 and S10.

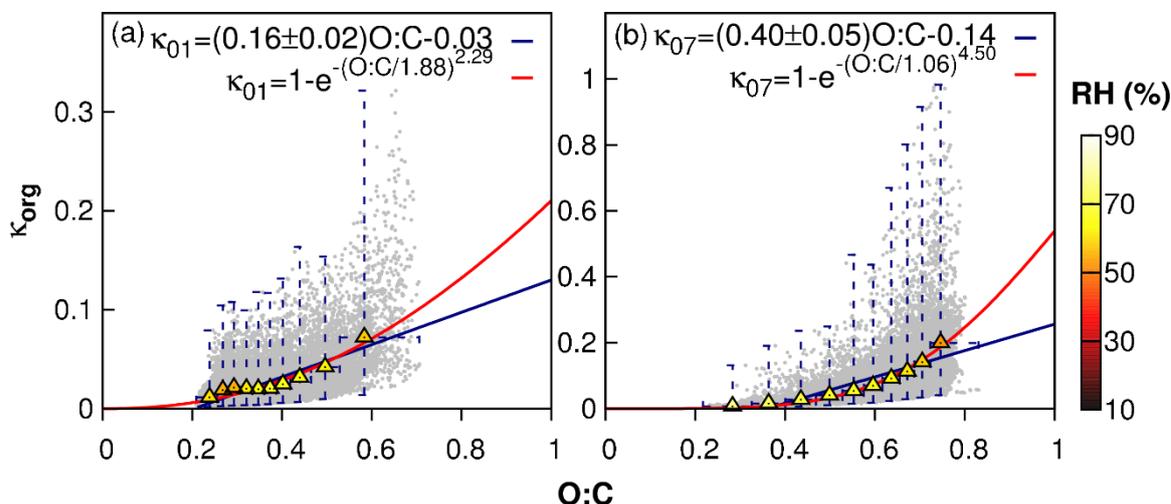


Figure 4. The correlation of hygroscopicity of organic aerosol (κ_{org}) and O:C ratio at 9 representative cities including Shenyang (SS), Beijing (BJ), Jinan (JN), Zhengzhou (ZZ), Xi'an (XA), Nanjing (NJ), Shanghai (SH), Chengdu (CD), and Guangzhou (GZ) in January (a) and July (b) of 2013. O:C ratios are categorized into 10 bins. In each bin, the ranges of O:C and κ_{org} are represented by bars.

The mean values of O:C and κ_{org} are represented by triangles colored by the averaged RH of each bin. The relationship between κ_{org} and O:C is fitted by a linear function with reduced major axis regression (blue lines) and an exponential function (red lines), respectively. κ_{01} and κ_{07} represent the fitted correlation for January and July, respectively.

Comment 11: Figure 7: The graph is not clear. In the main text, the authors discuss that the daily maximum of SOA occurs when RH is greater than 70% in both cities. The RH is higher than 70% all the time. The time of the day is up to 24 hours and not 25. The markers and boxes for (c) are not discussed whether they represent the left or right axis.

Response 11: We have removed this figure to avoid confusion.

Specific Comments

Comment 12: Line 34: Please elaborate more

Response 12: Now the text reads:

“However, the models typically assume that the organic particulate matter (OPM) is an ideal mixture and ignore the partitioning of water vapor to OPM.”

Comment 13: *Line 39: Please define generally with a statistical value that has meaning.*

Response 13: The text has been revised in L40-42 as follows:

“The modified model can generally capture the observed surface organic carbon (OC) with a correlation coefficient R of 0.7, and the surface OA with the mean fractional bias (MFB) and mean fractional error (MFE) of -0.28 and 0.54, respectively.”

Comment 14: *Line 91: Please elaborate more on the “purer condensed organics” for non-experts.*

Response 14: This should be condensed organics and has been revised accordingly.

Comment 15: *Line 99: “neglect” instead of “neglected”.*

Response 15: The text has been revised as instructed.

Comment 16: *Line 99: Please elaborate more on 1).*

Response 16: The text has been revised in L98-100 as following:

“1) the molecular structures and interactions of functional groups (-OH, -C=O, -COOH, etc.) of condensed organics (non-ideality);”

Comment 17: *Line 109: The sentence is missing a verb.*

Response 17: The sentence has been revised in L109-111 as following:

“Laboratory and field studies have observed water absorbed by SOA from a variety of precursor VOCs (Lambe et al., 2011; Zhao et al., 2016b; Asa-Awuku et al., 2010; Varutbangkul et al., 2006).”

Comment 18: *Line 128-131: The sentence is hard to read. Please rephrase.*

Response 18: The sentence has been revised in L126-129 as follows:

“Using UNiversal Functional Activity Coefficient (UNIFAC) method (Fredenslund et al., 1975) for calculating activity coefficients of the organic-water mixture, it was found that in the eastern U.S., where biogenic SOA dominated the OA, considering ALW_{org} leads to a

significant increase in predicted SOA (Pankow et al., 2015; Jathar et al., 2016).”

Comment 19: *Line 138: Which region?*

Response 19: The region refers to China. The text has been revised in L134-135 as follows:

“Previous modeling studies in China indicate that SOA was underpredicted (Lin et al., 2016; Jiang et al., 2012)”

Comment 20: *Line 163: (1) Don't change a line. (2) Also, why acidic conditions? Please, elaborate more. (3) Why is the reactive uptake of dicarbonyls, IEPOX and MAE in the “non-volatile” category?*

Response 20: (1) This is a mistake due to file format conversion and has been corrected. (2) This is a typo. SOA formed by isoprene oxidation under acidic conditions refers to IEPOX and MAE SOA based on chamber experiments (Lal et al., 2012; Lin et al., 2012; 2013). We have removed “isoprene oxidation under acidic conditions” in the revised text. (3) In the current model, we assume that the reactive uptake of dicarbonyls, IEPOX and MAE is irreversible, as an upper-limit estimation of SOA from these precursors. That’s why they are classified as non-volatile SOA. The text has been revised in L165-173 as following:

“SOA from dicarbonyls, IEPOX, and MAE were formed by irreversible reactive uptake and categorized as NV-SOA in the current model as well. Some studies investigated SOA from glyoxal, methylglyoxal, and IEPOX using detailed reactions and reversible pathways in models or observed as reversible processes in chamber experiments, leading to a relatively lower SOA yield compared to the surface-controlled irreversible uptake (Lim et al., 2013; Knote et al., 2014; Galloway et al., 2009; El-Sayed et al., 2018; Budisulistiorini et al., 2017). The non-volatile assumption used in this paper allows an upper-limit estimation of the importance of these additional SOA formation pathways.”

Comment 21: *Line 164: change to “was mostly”*

Response 21: This sentence has been moved to L162 in the revised manuscript and now it reads as follows:

“The SOA module mostly follows Pankow et al. (2015).”

Comment 22: *Line 165: Please elaborate more on the non-ideality calculation of the organic-water mixture for non-experts.*

Response 22: This sentence has been moved to L199-200 in the revised manuscript and now it reads as follows:

“POA is also involved in the calculation of activity coefficients for the

organic-water mixture.”

Comment 23: *Line 170: Is this the absorbing organic phase?*

Response 23: Yes. The text has been revised to “the absorbing organic phase”.

Comment 24: *Line 196: Change to “as water condenses”.*

Response 24: The text has been revised as instructed.

Comment 25: *Line 200: Please elaborate more on the “Kelvin effect neglected” for non-experts.*

Response 25: The text has been revised in L210-211 as following:

“Based on the κ -Köhler theory with linearly additive hygroscopic behavior of each component of the mixed particle”

Comment 26: *Line 204: Change to “can be estimated”.*

Response 26: The text has been revised as instructed.

Comment 27: *Line 242: Observations in 8 sub-regions of the domain during which period?*

Response 27: The text has been revised in L278-281 as follows:

“Table 1 lists model statistics of mean observation (OBS), mean prediction (PRE), mean bias (MB), gross error (GE) and correlation coefficient (R) based on WRF and observations at monitoring sites located in 8 sub-regions of the domain (Figure S1) during January and July of 2013.”

Comment 28: *Line 251: No significant improvements observed when applying the above additions means that the model is still missing a significant pathway to SOA formation, especially since OA in both seasons are dominated by SOA based on observations. This should be discussed in detail and in the context of previous studies and findings from AMS measurements in China.*

Response 28: We have revised the text and added a discussion about the underestimation of SOA in the current model. Please refer to Response 2 for more details.

Comment 29: *Line 258: Does it capture the observed diurnal variation? What is the R^2 or R of the timeseries of the modeled to observed values? What is the ratio of the two? In many cases, it seems that the difference is higher than a factor of 4. Is that a usual discrepancy? If so, how is much is it improved when incorporating the detailed SOA models?*

Response 29: Our model can capture the diurnal variation. The mismatching

of several peak values might be due to uncertainties in the emission inventory and the underestimate of SOA in the current model. The R of the modeled to observed OA is 0.55. The ratio of the averaged prediction to observation is 0.75. Since the model predicts a very small ratio of SOA to POA, the improvement from the detailed SOA model is insignificant. The small SOA/POA ratio might be due to the missing SOA from other pathways including POA aging and oxidation from IVOCs and VOC oxidation products. This has been explained in Response 2.

We revised the text in L301-304 to expand more discussion of the modeled and observed OA comparison:

“CMAQ can well capture the observed diurnal variation of OA in Beijing during wintertime, except for the underestimates of peak values. The correlation coefficient of modeled to observed OA is 0.55. We find a 25% underestimate of OA on average.”

Comment 30: Line 259: It would be great if “better” was described with statistical terminology. A way to describe the data and the comparison to modeled values would be to generate box and whiskers of the ratio of observations to modeled values for non-polluted and polluted days, respectively.

Response 30: We did a mistake in the MFB and MFE calculation for OA, which should be -0.28 and 0.54, respectively. We also calculated the biases on polluted and non-polluted days of OA. MFB and MFE of polluted days are -0.38 and 0.64, which are -0.26 and 0.52 for non-polluted days. The text has been revised in L306-309 as follows:

“The mean fractional bias (MFB) and mean fractional error (MFE) of polluted days are -0.38 and 0.64, respectively, which are worse than that of the non-polluted days (-0.26 for MFB and 0.52 for MFE). The overall MFB and MFE of OA during January are -0.28 and 0.54, within the criteria (MFB \leq ±0.6; MFE \leq 0.75) suggested by EPA (2007).”

Comment 31: Line 262-263: POA is not the primary contributor to OA in Beijing in winter. Many studies show that SOA is the major contributor and the path towards SOA formation is currently unknown and strongly dependent on LWC in the particles. Aging of POA not treated in the model is not guaranteed to be the main source of SOA.

Response 31: Please refer to Response 2.

Comment 32: Line 280: Here only one season is provided in terms of timeseries comparison of the model and obs. Please provide both seasons.

Response 32: Unfortunately, detailed chemical composition measurements for

aerosols are very limited in China during 2013. We only have observations of OC and OA in January of 2013 and PM_{2.5} in July of 2013 available for model evaluation.

Comment 33: *Line 285: Figure S5 shows the anthropogenic SOA and not the dicarbonyl SOA. Please separate the contributions and discuss them in the main text. Identifying the contribution of different compounds to SOA formation in China would be of great interest to the scientific community.*

Response 33: The contribution of each precursor to SOA of this episode has been shown in Hu et al. (2017) and will not be discussed in detail in the current study. The text has been revised in L344-346 as following:

“Anthropogenic emissions are the major sources of SOA (Figure S6), such as dicarbonyl products from the oxidation of xylene and toluene (Hu et al., 2017).

Comment 34: *Line 312: What about particulate nitrate?*

Response 34: The interactions between water-inorganics and water-organics are treated separately in the current model. We only focus on the water-organic interaction in the current study.

Comment 35: *Line 380: RH is higher than 70% all the time. What is the meaning of this sentence?*

Response 35: We have removed this figure to avoid confusion.

References:

- Hodzic, A., Jimenez, J. L., Madronich, S., Canagaratna, M. R., DeCarlo, P. F., Kleinman, L., and Fast, J.: Modeling organic aerosols in a megacity: potential contribution of semi-volatile and intermediate volatility primary organic compounds to secondary organic aerosol formation, *Atmos. Chem. Phys.*, 10, 5491-5514, 10.5194/acp-10-5491-2010, 2010.
- Hu, J., Chen, J., Ying, Q., and Zhang, H.: One-year simulation of ozone and particulate matter in China using WRF/CMAQ modeling system, *Atmos. Chem. Phys.*, 16, 10333-10350, 10.5194/acp-16-10333-2016, 2016.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of Organic Aerosols in the Atmosphere, *Science*, 326, 1525-1529, 10.1126/science.1180353, 2009.
- Kota, S. H., Schade, G., Estes, M., Boyer, D., and Ying, Q.: Evaluation of MEGAN predicted biogenic isoprene emissions at urban locations in Southeast Texas, *Atmospheric Environment*, 110, 54-64, 2015.
- Lal, V., Khalizov, A.F., Lin, Y., Galvan, M.D., Connell, B.T., Zhang, R., 2012. Heterogeneous reactions of epoxides in acidic media. *J. Phys. Chem. A* 116, 6078e6090.
- Li, J., Cleveland, M., Ziemba, L. D., Griffin, R. J., Barsanti, K. C., Pankow, J. F., and Ying, Q.: Modeling regional secondary organic aerosol using the Master Chemical Mechanism, *Atmos. Environ.*, 102, 52-61, <https://doi.org/10.1016/j.atmosenv.2014.11.054>, 2015.
- Lin, Y.H., Zhang, Z.F., Docherty, K.S., Zhang, H.F., Budisulistiorini, S.H., Rubitschun, C.L., Shaw, S.L., Knipping, E.M., Edgerton, E.S., Kleindienst, T.E., Gold, A., Surratt, J.D., 2012b. Isoprene epoxydiols as precursors to secondary organic aerosol formation: acid-catalyzed reactive uptake studies with authentic compounds. *Environ. Sci. Technol.* 46, 250-258.
- Lin, Y.-H., Zhang, H., Pye, H. O. T., Zhang, Z., Marth, W. J., Park, S., Arashiro, M., Cui, T., Budisulistiorini, S. H., Sexton, K. G., Vizuete, W., Xie, Y., Luecken, D. J., Piletic, I. R., Edney, E. O., Bartolotti, L. J., Gold, A., and Surratt, J. D.: Epoxide as a precursor to secondary organic aerosol formation from isoprene photo-oxidation in the presence of nitrogen oxides, *P. Natl. Acad. Sci. USA*, 110, 6718-6723, doi:10.1073/pnas.1221150110, 2013.
- Liu, J., Shen, J., Cheng, Z., Wang, P., Ying, Q., Zhao, Q., Zhang, Y., Zhao, Y., and Fu, Q.: Source apportionment and regional transport of anthropogenic secondary organic aerosol during winter pollution periods in the Yangtze River Delta, China, *Sci. Total Environ.*, 710, 135620, <https://doi.org/10.1016/j.scitotenv.2019.135620>, 2020.
- Murphy, B. N., Woody, M. C., Jimenez, J. L., Carlton, A. M. G., Hayes, P. L., Liu, S., Ng, N. L., Russell, L. M., Setyan, A., Xu, L., Young, J., Zaveri, R. A., Zhang, Q., and Pye, H. O. T.: Semivolatile POA and parameterized total combustion SOA in CMAQv5.2: impacts on source strength and partitioning, *Atmos. Chem. Phys.*, 17, 11107-11133, 10.5194/acp-17-11107-2017, 2017.

Pankow, J. F., Marks, M. C., Barsanti, K. C., Mahmud, A., Asher, W. E., Li, J., Ying, Q., Jathar, S. H., and Kleeman, M. J.: Molecular view modeling of atmospheric organic particulate matter: Incorporating molecular structure and co-condensation of water, *Atmospheric Environment*, 122, 400-408, <http://dx.doi.org/10.1016/j.atmosenv.2015.10.001>, 2015.

Pye, H. O. T., Murphy, B. N., Xu, L., Ng, N. L., Carlton, A. G., Guo, H., Weber, R., Vasilakos, P., Appel, K. W., Budisulistiorini, S. H., Surratt, J. D., Nenes, A., Hu, W., Jimenez, J. L., Isaacman-VanWertz, G., Misztal, P. K., and Goldstein, A. H.: On the implications of aerosol liquid water and phase separation for organic aerosol mass, *Atmos. Chem. Phys.*, 17, 343-369, 10.5194/acp-17-343-2017, 2017.

Qiao, X., Ying, Q., Li, X., Zhang, H., Hu, J., Tang, Y., and Chen, X.: Source apportionment of PM_{2.5} for 25 Chinese provincial capitals and municipalities using a source-oriented Community Multiscale Air Quality model, *Sci. Total Environ.*, 612, 462-471, <https://doi.org/10.1016/j.scitotenv.2017.08.272>, 2018.

Robinson, A. L., Donahue, N. M., Shrivastava, M. K., Weitkamp, E. A., Sage, A. M., Grieshop, A. P., Lane, T. E., Pierce, J. R., and Pandis, S. N.: Rethinking Organic Aerosols: Semivolatile Emissions and Photochemical Aging, *Science*, 315, 1259-1262, 10.1126/science.1133061, 2007.

Shi, Z., Li, J., Huang, L., Wang, P., Wu, L., Ying, Q., Zhang, H., Lu, L., Liu, X., Liao, H., and Hu, J.: Source apportionment of fine particulate matter in China in 2013 using a source-oriented chemical transport model, *Science of the Total Environment*, 601, 1476-1487, <http://dx.doi.org/10.1016/j.scitotenv.2017.06.019>, 2017.

Shrivastava, M. K., Lane, T. E., Donahue, N. M., Pandis, S. N., and Robinson, A. L.: Effects of gas particle partitioning and aging of primary emissions on urban and regional organic aerosol concentrations, 113, 10.1029/2007jd009735, 2008.

Situ, S., Wang, X., Guenther, A., Zhang, Y., Wang, X., Huang, M., Fan, Q., and Xiong, Z.: Uncertainties of isoprene emissions in the MEGAN model estimated for a coniferous and broad-leaved mixed forest in Southern China, *Atmos. Environ.*, 98, 105-110, <https://doi.org/10.1016/j.atmosenv.2014.08.023>, 2014.

Sun, Y., Du, W., Fu, P., Wang, Q., Li, J., Ge, X., Zhang, Q., Zhu, C., Ren, L., Xu, W., Zhao, J., Han, T., Worsnop, D. R., and Wang, Z.: Primary and secondary aerosols in Beijing in winter: sources, variations and processes, *Atmos. Chem. Phys.*, 16, 8309-8329, 10.5194/acp-16-8309-2016, 2016.

Sun, Y. L., Wang, Z. F., Fu, P. Q., Yang, T., Jiang, Q., Dong, H. B., Li, J., and Jia, J. J.: Aerosol composition, sources and processes during wintertime in Beijing, China, *Atmos. Chem. Phys.*, 13, 4577-4592, 10.5194/acp-13-4577-2013, 2013.

Tkacik, D. S., Presto, A. A., Donahue, N. M., and Robinson, A. L.: Secondary Organic Aerosol Formation from Intermediate-Volatility Organic Compounds: Cyclic, Linear, and Branched Alkanes, *Environmental Science & Technology*, 46, 8773-8781, 10.1021/es301112c, 2012.

Wang, P., Schade, G., Estes, M., and Ying, Q.: Improved MEGAN predictions of biogenic isoprene in the contiguous United States, *Atmospheric Environment*, 148, 337-351, <http://dx.doi.org/10.1016/j.atmosenv.2016.11.006>, 2017.

Ying, Q., Cureño, I. V., Chen, G., Ali, S., Zhang, H., Malloy, M., Bravo, H. A., and Sosa, R.: Impacts of Stabilized Criegee Intermediates, surface uptake processes and higher aromatic secondary organic aerosol yields on predicted PM_{2.5} concentrations in the Mexico City Metropolitan Zone, *Atmos. Environ.*, 94, 438-447, <https://doi.org/10.1016/j.atmosenv.2014.05.056>, 2014.

Ying, Q., Li, J., and Kota, S. H.: Significant Contributions of Isoprene to Summertime Secondary Organic Aerosol in Eastern United States, *Environ. Sci. Technol.*, 49, 7834-7842,

10.1021/acs.est.5b02514, 2015.

Zhao, B., Wang, S., Donahue, N. M., Jathar, S. H., Huang, X., Wu, W., Hao, J., and Robinson, A. L.: Quantifying the effect of organic aerosol aging and intermediate-volatility emissions on regional-scale aerosol pollution in China, *Sci. Rep.*, 6, 28815, 10.1038/srep28815, 2016.

Zhao, J., Qiu, Y., Zhou, W., Xu, W., Wang, J., Zhang, Y., Li, L., Xie, C., Wang, Q., Du, W., Worsnop, D. R., Canagaratna, M. R., Zhou, L., Ge, X., Fu, P., Li, J., Wang, Z., Donahue, N. M., and Sun, Y.: Organic Aerosol Processing During Winter Severe Haze Episodes in Beijing, *J. Geophys. Res.*, 124, 10248-10263, 10.1029/2019jd030832, 2019.

Zhao, Y., Hennigan, C. J., May, A. A., Tkacik, D. S., de Gouw, J. A., Gilman, J. B., Kuster, W. C., Borbon, A., and Robinson, A. L.: Intermediate-Volatility Organic Compounds: A Large Source of Secondary Organic Aerosol, *Environ. Sci. Technol.*, 48, 13743-13750, 10.1021/es5035188, 2014.

Zheng, B., Zhang, Q., Zhang, Y., He, K. B., Wang, K., Zheng, G. J., Duan, F. K., Ma, Y. L., and Kimoto, T.: Heterogeneous chemistry: a mechanism missing in current models to explain secondary inorganic aerosol formation during the January 2013 haze episode in North China, *Atmos. Chem. Phys.*, 15, 2031-2049, 10.5194/acp-15-2031-2015, 2015.

Response to Referee #2

The authors thank the reviewer for the comments that improve the quality of the paper. The detailed responses are given as follows. The reviewer comments are shown in italic fonts, the responses are in regular font, and the revised text is in bold font.

In this paper, Li et al. have modified the CMAQ model to take into account the impacts of water partitioning and polarity of organic compounds on SOA formation. The model was applied over Eastern China to estimate the regional and seasonal impacts of these modifications on SOA and the aerosol water content. This study may have potential to contribute in the organic aerosol modeling field but major revisions needs to be done before publication. In particular, I have several concerns regarding the validity of the scientific methodology used and the presentation of the study. Therefore, I would recommend publication only if these comments will be addressed and fundamental changes will be contacted.

Major comments

Comment 1: *Page 4 line 94: The majority of current CTMs have replaced the 2-product model with the VBS approach. Please make this clear and refer to the 2-product model of Odum et al. (1996) for historical reasons.*

Response 1: The text has been revised in the manuscript (L88-95) to make this clear:

“The formation of condensed organic products is commonly represented by lumped surrogate SVOCs in a 2-product model with volatilities and SVOC yields fitted to chamber experiments (Odum et al., 1996). To better represent the volatility of primary organic aerosol (POA) and the multi-generation oxidation of SVOCs to a wider range, Donahue et al. (2006) proposed the volatility basis set (VBS) model in which the mass yields of SVOCs are fitted to a fixed number of volatility bins (usually 0.01-105 $\mu\text{g m}^{-3}$). The VBS model has been adopted by several CTMs (such as WRF-Chem, GEOS-Chem, etc.).”

Comment 2: *Page 6, line 162: Please add the appropriate references to support the nonvolatile nature of the products by these oxidation pathways.*

Response 2: Formation of these non-volatile SOA was traditionally treated in the CMAQ model, except for dicarbonyls, IEPOX, and MAE that are assumed to form SOA by irreversible reactive uptake in our model, as an upper-limit estimation of SOA from these precursors. The text has been revised to make this clear (L162-173 in the manuscript):

“The SOA module mostly follows Pankow et al. (2015). Two types of SOA as traditionally treated in CMAQ were considered, “semi-volatile” (SV) portion that formed via equilibrium absorption-partitioning of SVOCs, and

“non-volatile” (NV) portion that includes the oligomers and SOA formed via direct oxidation of aromatics at low-NO_x. SOA from dicarbonyls, IEPOX, and MAE were formed by irreversible reactive uptake and categorized as NV-SOA in the current model as well. Some studies investigated SOA from glyoxal, methylglyoxal, and IEPOX using detailed reactions and reversible pathways in models or observed as reversible processes in chamber experiments, leading to a relatively lower SOA yield compared to the surface-controlled irreversible uptake (Lim et al., 2013; Knote et al., 2014; Galloway et al., 2009; El-Sayed et al., 2018; Budisulistiorini et al., 2017). The non-volatile assumption used in this paper allows an upper-limit estimation of the importance of these additional SOA formation pathways.”

Comment 3: *Page 6 line 149: There is no discussion in the methodology about the observations and the statistical analysis metrics used to evaluate the model performance. Especially for the OA observations, there is no reference provided or description of the methods used.*

Response 3: Thank you for pointing this out. We have included details of observation data, statistical analysis metrics for both meteorology and aerosols in the revised manuscript:

In L265-276

“The meteorological inputs and emissions have been used in several previous publications. Model performance on meteorological parameters (temperature and RH), gaseous species and gas and aerosol concentrations have been extensively evaluated (Hu et al., 2016; Hu et al., 2017; Qiao et al., 2018; Shi et al., 2017). A summary of the model performance related to this study is provided below. Observed meteorological data were obtained from the National Climatic Data Center (<ftp://ftp.ncdc.noaa.gov/pub/data/noaa>). Observations of OC at two urban locations, Beijing (Cao et al., 2014; Wang et al., 2015) and Guangzhou (Lai et al., 2016) and OA in Beijing (Sun et al., 2014) during January of 2013 as well as surface PM_{2.5} at several monitoring sites during July of 2013 from China National Environmental Monitoring Center (<http://113.108.142.147:20035/emcpublish/>) were used to evaluate model estimates of aerosols. Details of measurement methodology and uncertainties of observations are listed in the corresponding references.”

In L278-282

“Table 1 lists model statistics of mean observation (OBS), mean prediction (PRE), mean bias (MB), gross error (GE) and correlation coefficient (R) based on WRF and observations at monitoring sites located in 8 sub-regions of the domain (Figure S1) during January and July of 2013. The benchmarks for the MM5 model (another meteorology model) of 4-12km horizontal resolution suggested by Emery et al. (2001) are also listed in

the table.”

In L294-296

“Overall, the mean fractional bias (MFB) and mean fractional error (MFE) of OC are -0.20 and 0.27, within the criteria ($MFB \leq \pm 0.6$; $MFE \leq 0.75$) suggested by EPA (2007).”

In L308-309:

“The overall MFB and MFE of OA during January are -0.28 and 0.54, within the criteria ($MFB \leq \pm 0.6$; $MFE \leq 0.75$) suggested by EPA (2007).”

Comment 4: *Page 7 line 179: More details are needed here. (1) How the model defines the low NO_x and high NO_x conditions? Which compounds each of the lump species represent? What is the difference between the lumped species of the same precursor (e.g., BNZ1, BNZ2, BNZ3)? (2) Can you include the aerosol yields for each lumped species in tables S1 and S2? Are these aerosol yields NO_x-dependent? (3) What does the SVP stands for in Tables S1 and S2?*

Response 4: (1) SOA formation in CMAQv5.0.2 is based on the frame of a previous version 4.7.1. All the details about “high” and “low” NO_x conditions (based on chamber experiments of corresponding VOCs), lumping species and method of each precursor, and the yields of precursors from parent VOCs have been documented by Carlton et al. (2010) and summarized in the revised supplemental materials as following:

“The CMAQ model treats high and low NO_x SOA formation pathways during OH oxidation by allowing the lumped RO₂ radical to competitively react with HO₂ and NO. Using the lumped ARO1 species as an example, an SOA formation specific RO₂ radical ARO1RO₂ is added as a gas phase reaction product with OH:



The ARO1RO₂ can react with both HO₂ and NO, as shown in the following two reactions:



Details of the determination of the rate constants can be found in Carlton et al. (2010). The TOLNRXN and TOLHRXN are counter species that track how much ARO1 is reacted through low NO_x and high NO_x pathways, respectively, in one gas chemistry time step. The concentrations of these counter species are passed into the aerosol module to calculate the formation semi-volatile products (TOL1 and TOL2) in the high NO_x pathway and non-volatile products (TOL3) in the low NO_x pathway, using the mass-specific yields, as listed in Table S1 and S2. Equilibrium partitioning of TOL1 and TOL2 in the gas phase and their counterparts ATOL1 and ATOL2 in the organic phase are affected by temperature and the amount of absorbing organics in the aerosol phase. Similar

treatments are applied to the other lumped aromatic compounds ARO2, with xylene as a representative and most abundant species in that group, and to benzene. SOA formation from lumped long-chain alkene species ALK5, and isoprene and monoterpenes are not considered as NO_x dependent and are represented by equilibrium partitioning of one or two semi-volatile oxidation products. Details of the mass-specific yields of semi-volatile products and other related parameters can be found in Table S1 and S2.”

We revised the text in (L188-190) to make it clear:

“More details about the lumped precursors such as formation conditions (“high” and “low” NO_x), lumping species and method, and yields from parent VOCs can be found in Carlton et al. (2010) and summarized in SI.”

(2) In the CMAQ model, the amount of SOA can form after a precursor reacts with OH, O₃ or NO₃ depends on the volatility of the products, which is temperature dependent, and the amount of the absorbing organics. The mass yields of the semi-volatile or non-volatile products are included in Table S1 and S2 in the revised supplementary materials. For more details, we refer the readers to Carlton et al. (2010) and the references therein.

(3) SVP is the saturation vapor pressure. We have explained this in the corresponding tables.

Comment 5: *Page 8, lines 194-198: (1) Does the absorbing phase of equation 1 includes only the water associated with the organics (from eq. 3) or it includes the total water (including the water associated with the inorganic aerosol components)? (2) Under high RH (higher than the organic/inorganic phase separation RH, SRH), the aerosol organic phase is well mixed with the inorganic salts and, therefore, the aerosol water associated with the inorganic constituents can also contribute to the SOA absorbing medium (Pye et al., 2017). Please clarify what you have assumed here and add the relative discussion.*

Response 5: (1) The absorbing phase of equation 1 only includes organic and water associated with organics when considering water-organic interactions. We explained this in the original text L195-196:

“In addition to organic compounds, water partitioning into OPM is enabled according to Eq 1 and Eq 2. In such a case, the absorbing phase in Eq 1 includes both organic aerosols and water partitioning into OPM.”

(2) We didn’t consider the mixing of organic and inorganic phase in this study and assumed that they are always two distinct aerosol phases without direct interactions. The phase separation RH (SRH) depends on the OM/OC ratio of the organic phase. Unlike the conditions modeled by Pye et al. (2017) for the

southeast US where SOA is often the dominant OA component, the winter episode we modeled is dominated by primary emitted organic aerosols thus with a relatively low OM/OC (~1.4-1.6). The SRH based on equation (7) of Pye et al. (2017) is ~97%-99%. The summer episode has more contributions of SOA to OA, with OM/OC~1.8, which corresponds to an SRH of 87%. Thus, we don't expect interactions of organic and inorganic phases to occur in high frequency to greatly influence the model results. We assumed no interactions between inorganic and organic phases in the current model. We have also revised text in L207-208:

"In the current model, we assumed no interactions between the inorganic and organic phases."

Comment 6: Page 8, Equation 3: This equation gives the volume of water associated with the organic fraction of the aerosol. However, ALW on the left-hand side refers to the mass of water. Please correct.

Response 6: The equation has been corrected in the revised text L212-214:

"

$$ALW_{org} = \rho_w V_{org} \kappa_{org} \frac{a_w}{1 - a_w} \quad (\text{Eq3})$$

where ρ_w is the density of water (assumed to be 1 g cm⁻³), V_{org} is the volume concentration of organics, and a_w is the water activity (assumed to be the same as RH)."

Comment 7: Page 8, lines 203-204: How do you estimate the hygroscopicity?

Response 7: We used Eq3 to estimate the hygroscopicity. This has been clarified in the revised text in L214-216:

"Since ALW_{org} in this study is calculated mechanistically using the partitioning theory, κ_{org} can be estimated by rearranging Eq3:

$$\kappa_{org} = \frac{ALW_{org}}{\rho_w V_{org}} \times \frac{1 - a_w}{a_w} \quad (\text{Eq4})$$

"

Comment 8: Page 8, lines 204-205: How do you calculate the ALW? Are you using κ_{org} and the eq3? Do you use the kappa hygroscopicity to calculate the ALW? If so, how you estimate the kappa?

Response 8: In this study, the ALW_{org} is independently calculated by mechanistically allowing water molecules to partition into the organic phase with UNIFAC calculated activity. In such a case, Eq 3 can be used to provide an independent estimation of κ_{org} . Linear regression analysis can be performed using the calculated κ_{org} against the model calculated O/C ratios, as shown in Figure 4. We have removed this sentence to avoid confusion.

Comment 9: *Page 9, lines 212-216: Add a reference to tables S1 and S2. Furthermore, why the values of OM:OC in tables S1/S2 are different than the values provided by your reference (Pye et al., 2017)?*

Response 9: We used OM:OC ratio in Pankow et al. (2015). There were mistakes in the original Table S1 and S2 and have been corrected in the revised supplemental materials. The text has also been revised in L228-229 as following:

“The OM:OC ratio of each SOA component follows Pankow et al. (2015) as shown in Table S1-S2.”

Comment 10: *Page 9 lines 217-218: Can you add the total size of the model domain?*

Response 10: The text has been revised in L231-233 as follows:

“The simulation domain has a horizontal resolution 36 km × 36 km (100 × 100 grids) and a vertical structure of 18 layers up to 21 km, which covers eastern China as shown in Figure S1.”

Comment 11: *Page 9 lines 233: (1) What boundary conditions are used? Please make a comment on how these can affect the simulation results. (2) I would recommend adding spatial maps of primary organic aerosol emissions and SOA precursor emissions and summarizing in a Table the domain average emission rates of POA and each SOA precursors.*

Response 11: (1) We used a predefined boundary profile in CMAQ that represents a clean continental condition. The northern and western boundaries, as well as areas to the further north and west, are mostly remote areas with much lower emissions. The mountains in the north and west part of the domain also limit the influence of emissions enter from the boundaries to the central part of the domain. The influence of marine air from the south and east boundaries is also small, as local emissions dominate the concentrations.

(2) A figure and a table showing emissions of POA and SOA precursors were added to the revised supplemental materials as Figure S2 and Table S4.

Comment 12: *Pages 9-10 lines 234-238: (1) This paragraph needs to be expanded and written in a separate section. In this section the authors should describe in more detail the following: i) Basecase simulation. Please explain how the default CMAQ is simulating POA and SOA and how different is this modelling configuration with the one the authors are testing, ii) Sensitivity simulations. Please explain in much more detail the sensitivity simulations conducted in this work. (2) In addition the authors say that they have conducted three sensitivity scenarios named S1, S2 and S3. In their manuscript they only show results from S3 and they never discuss the results of S1 and S2.*

Response 12: (1) Details of how default CMAQ simulates SOA formation through equilibrium partitioning of lumped semi-volatile products into the

organic phase (which include both POA and SOA) have been described by Carlton et al. (2010) and Hu et al. (2017) so we don't think it is necessary to repeat it here. We expanded each simulation scenario in details in the revised text in L249-259:

“Four scenarios are investigated in this study. The base case (BS) applies the default secondary organic aerosol module of CMAQ v5.0.1. In this case, no water partitioning into OPM is considered. Lumped semi-volatile products from the oxidation of various precursors partition into a single organic phase, which is considered as an ideal mixture of POA and SOA with $\gamma_{org}=1$. The water case (S1) includes water partitioning into OPM, which is again considered as an ideal solution ($\gamma_{org}=1$ and $\gamma_{H_2O}=1$). The UNIFAC case (S2) considers the interaction between organic constituents with UNIFAC calculated activity coefficients ($\gamma_{org}\neq 1$) but does not allow water partitioning into OPM. The combined case (S3) allows both water partitioning and interactions between all constituents (including water and organics) using UNIFAC calculated activity coefficients ($\gamma_{org}\neq 1$ and $\gamma_{H_2O}\neq 1$). ”

(2) The impacts of S1 and S2 were discussed in Section 4 of the original manuscript Page 16 L405-419. We also included a description of the outlines of the results and discussions in the original manuscript Page 6 L140-148. The text has been revised in L259-262 to make it clear:

“The results of BS and S3 are used to examine the overall impacts of water partitioning into OPM and polarity of organics on SOA and ALWorg, as shown in Section 3.1-3.4. The separate influences of those two processes on SOA from S1 and S2 are discussed in Section 3.5.”

Comment 13: Page 10 line 240: The section “Model evaluation” is extremely problematic and raises questions on the validity of the modelling results given that the model evaluation is insufficient. More specifically: 1) Given that the CMAQ default configuration has been modified to consider the importance of water and organic compound polarity on SOA formation, an accurate evaluation of the updated model performance is needed. 2) You should compare the model results for organic compounds during both July and January 2013. Currently, the evaluation includes a comparison with OC observations only during January over only three locations of the relatively large model domain. Furthermore, please mention in the text what factor have you used to convert the modeled OA to OC. 3) The total PM2.5 measurements have been used to evaluate the model performance during July without explaining the rationale of this choice since the focus of this study is solely the organic fraction of the aerosols. I suggest removing the PM2.5 evaluation or at least moving it to the supplement. 4) Can you include more OC/OA observations over other areas of their domain in your evaluation? 5) It is also important to compare the simulated

POA and SOA against observations (e.g., from AMS). Furthermore, it would be helpful to show how the model performance against SOA measurement changes between the BC and the S1, S2, S3 cases.

Response 13: (1) For the updates in CMAQ except for water partition into OPM and non-ideality of the organic-water mixture, previous studies have extensively examined the model performances and will not be further discussed in detail in this work. The text has been revised in L149-161 to reflect this:

“Model configurations were largely based on that used by Hu et al. (2016) as summarized below. Firstly, SAPRC-11 was expanded for a more detailed treatment of isoprene oxidation and tracking dicarbonyl products (glyoxal and methylglyoxal) from different groups of major precursors (Ying et al., 2015). Secondly, SOA from isoprene epoxydiols (IEPOX), methacrylic acid epoxide (MAE) and dicarbonyls through surface-controlled irreversible reactive uptake were added (Hu et al., 2017; Li et al., 2015; Liu et al., 2020; Ying et al., 2015). Thirdly, the heterogeneous formation of secondary nitrate and sulfate from NO₂ and SO₂ reaction on the particle surfaces (Ying et al., 2014) were added, which is an important source of secondary inorganic aerosols (Zheng et al., 2015) and improves model estimates of nitrate and sulfate (Qiao et al., 2018; Shi et al., 2017). Fourthly, SOA yields were corrected for vapor wall loss (Zhang et al., 2014). Impacts of the above updates on model performances have been extensively discussed in the cited work and will not be further investigated in the current study.”

(2) Unfortunately, detailed chemical composition measurements for aerosols are very limited in China during 2013. We only have observations of OC and OA in January of 2013 and PM_{2.5} in July of 2013 available for model evaluation. Thus, we opt to use the most relevant data to provide a very limited assessment of the capability of the model in predicting SOA. The factors for OA to OC conversion follow the OM:OC ratio listed in Table S1 and S2 for SOA. POC is directly predicted by the model. The text has been revised in L291-292 to make this clear:

“The factors used to convert SOA to OC (SOC) are listed in Table S1-S2. OC from POA (POC) is directly predicted by the model.”

(3) Even though limited OC/OA measurements are available to us during this period, the base case model is later applied by another research group to model wintertime SOA formation in east China (Liu et al., 2020). The predicted OC and SOA agree well with observations (Figure 2 of Liu et al. 2020), and the model performance statistics for OC and SOA are similar to those of PM_{2.5}. We agree with the reviewer that PM_{2.5} is not an ideal indicator to evaluate the capability of the model in predicting SOA, however, as a significant fraction of PM_{2.5} in July is secondary, this still provides an indirect assessment of the model prediction of oxidation capacity of the atmosphere, which is important for

SOA formation. Additional modeling studies are needed to evaluate the performance of the model in summer.

We have explained this in the revised text (L327-329) as following:

“Due to the lack of observed OC and OA in July of 2013, as an alternative, model performances are evaluated by comparing predicted and observed PM_{2.5} at ground sites (Figure S1) as shown in Figure S3.”

(4) We do not have more OC and OA data for the simulating episode of 2013.

(5) We did not have SOA observations in this episode. We compared the modeled SOA/POA ratio with AMS observations from other literature, finding a significant underestimation in the current model. This bias might be due to missing SOA from combustion (intermediate volatile organic compounds, IVOCs) and not treating POA as semi-volatile. We added a discussion about the model bias in the revised manuscript in L309-326:

“Again, no apparent changes of SOA nor OA are observed between case S3 and BS (not shown), since POA is predicted to be the primary contributor to OA at Beijing in winter in the current model, with an averaged SOA/POA ratio of 0.12. This ratio is much lower than the field observation of about 0.45-1.94 (Zhao et al., 2019; Sun et al., 2013; Sun et al., 2016). The bias might be due to the missing SOA converted by partitioning and aging of semi-volatile POA as well as oxidation from intermediate volatile organic compounds (IVOCs) and VOC oxidation products. Those pathways are shown to be important for SOA formation by modeling, field and chamber studies (Hodzic et al., 2010; Jimenez et al., 2009; Murphy et al., 2017; Robinson et al., 2007; Shrivastava et al., 2008; Tkacik et al., 2012; Zhao et al., 2014; Zhao et al., 2016a).

A sensitivity test was performed by using the newest CMAQ model version 5.3.1 that includes all the above processes in the aerosol module. The SOA/POA ratio in Beijing is improved greatly to be 0.83 in winter. However, high uncertainties still exist in the emissions of the involved precursors and characterization of SOA formation through these processes, needing further constrains by observations. Their influences on water partitioning into OPM and non-ideality of the organic-water mixture on SOA will be evaluated in a future study.”

Since SOA is underestimated, no significant differences in BS and S3 are observed. Case S1 and S2 are designed for the sensitivity test of model results to water partitioning and non-ideality of condensed organics separately. Therefore, we did not evaluate model performances from S1 and S2.

Comment 14: *Page 10 lines 251-253: The authors state here that the impacts of water-cocondensation and polarity of organic condensed species on SOA*

formation are not significant during winter. This highlights the need to evaluate their model results during July where they have found significant changes with the basecase simulation. Furthermore, the results from the three sensitivity simulations should be evaluated individually.

Response 14: Unfortunately, we have no observations of OC, OA or (and) SOA of July 2013. We tried to evaluate model performances by comparing the predicted and observed PM_{2.5} as an alternative. This has been explained in the revised manuscript in L327-329:

“Due to the lack of observed OC and OA in July of 2013, as an alternative, model performances are evaluated by comparing predicted and observed PM_{2.5} at ground sites (Figure S1) as shown in Figure S3.”

Since SOA is underestimated, no significant differences in BS and S3 are observed. Case S1 and S2 are designed for the sensitivity test of model results to water partitioning and non-ideality of condensed organics separately. Therefore, we did not evaluate model performances from S1 and S2.

Comment 15: *Page 10 lines 254-256: These are indeed possible factors. Can the authors comment, based on their analysis, which of these two possible factors is more important and try to be more specific? A comparison with AMS observations would be helpful here.*

Response 15: We did a sensitivity test by simulating the same episode with the current CMAQv5.3.1 in which POA was treated as semi-volatile and aging in the gas phase. Also, a missing source of SOA from intermediate VOCs (IVOCs) oxidation and aging of IVOCs and VOCs oxidation products (pcSOA) was added in CMAQv5.3.1. The results showed that the modeled SOA/POA has been improved from 0.12 of case S3 to 0.83 (Figure R1), more close to AMS observations. However, there were still some peak values underestimated by the model, which might be due to the uncertainties of POA emissions. We have revised the text (L309-326), as mentioned in Response 13 (5).

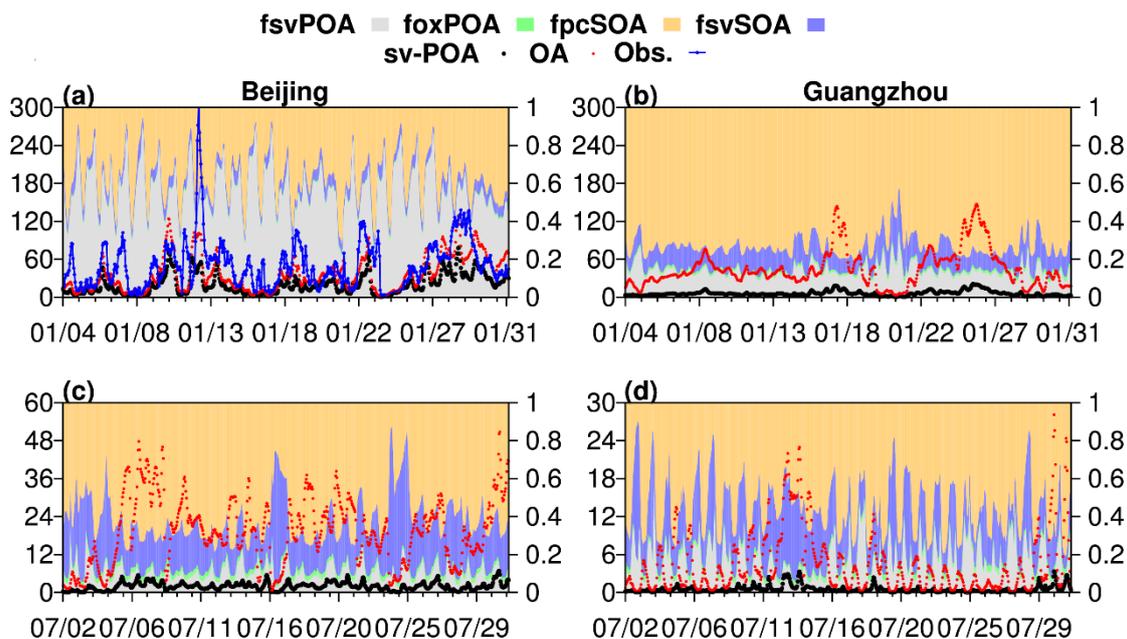


Figure R1. Modeled concentration of semi-volatile POA (sv-POA), OA and fraction of each organic aerosol component fsvPOA (sv-POA), foxPOA (oxidation of sv-POA), fpcSOA (pc-SOA) and fsvSOA (traditional SOA) in Beijing (a, c) and Guangzhou (b, d). Observations of OA in January 2013 at Beijing (Obs.) are also included in (a). The left axis is the concentration ($\mu\text{g m}^{-3}$) and the right axis is the fraction of OA components.

Comment 16: Page 11 line 263: *The aging of POA, under specific conditions can enhance the SOA formation, especially over polluted areas. Can the authors comment how this important omission of their model configuration can affect their result? Once again, a comparison against POA and SOA from AMS observations will be helpful to identify the limitations of their model due to the treatment of POA as non-volatile and non-reactive.*

Response 16: We did a sensitivity test by CMAQv5.3.1 that includes SOA from POA aging, IVOCs oxidation, and aging of IVOCs and VOCs oxidation products. The modeled SOA/POA is improved greatly in Beijing with no significant improvement in OA compared to the results of case S3. A significant improvement of SOA was observed from the contribution from IVOCs oxidation, and the aging of IVOCs and VOCs oxidation products (pc-SOA). Discussions about this have been included in the revised text in L309-326, as mentioned in Response 15.

Also, we examined the sensitivity of SOA and organic liquid water (ALW_{org}) to pcSOA and POA in an offline calculation in Beijing, Guangzhou, Jinan, and Nanjing. POA has the same properties as we used in the model. Non-volatile isoprene SOA is taken to represent pcSOA as their similarities in saturation vapor pressure and O:C ratio. We found that both SOA and ALW_{org} are

positively correlated with pcSOA, increased by 2-5 times in different locations when pcSOA increased by 2 times.

Comment 17: Page 11 line 264: Please provide two spatial maps of the fraction SOA/Total OA during January and July 2013 so as to show the contribution of POA and SOA to total OA during each simulation period.

Response 17: We have added a figure of SOA/OA ratio by case BS and S3 (as shown below) in the revised supplemental materials (Figure S13).

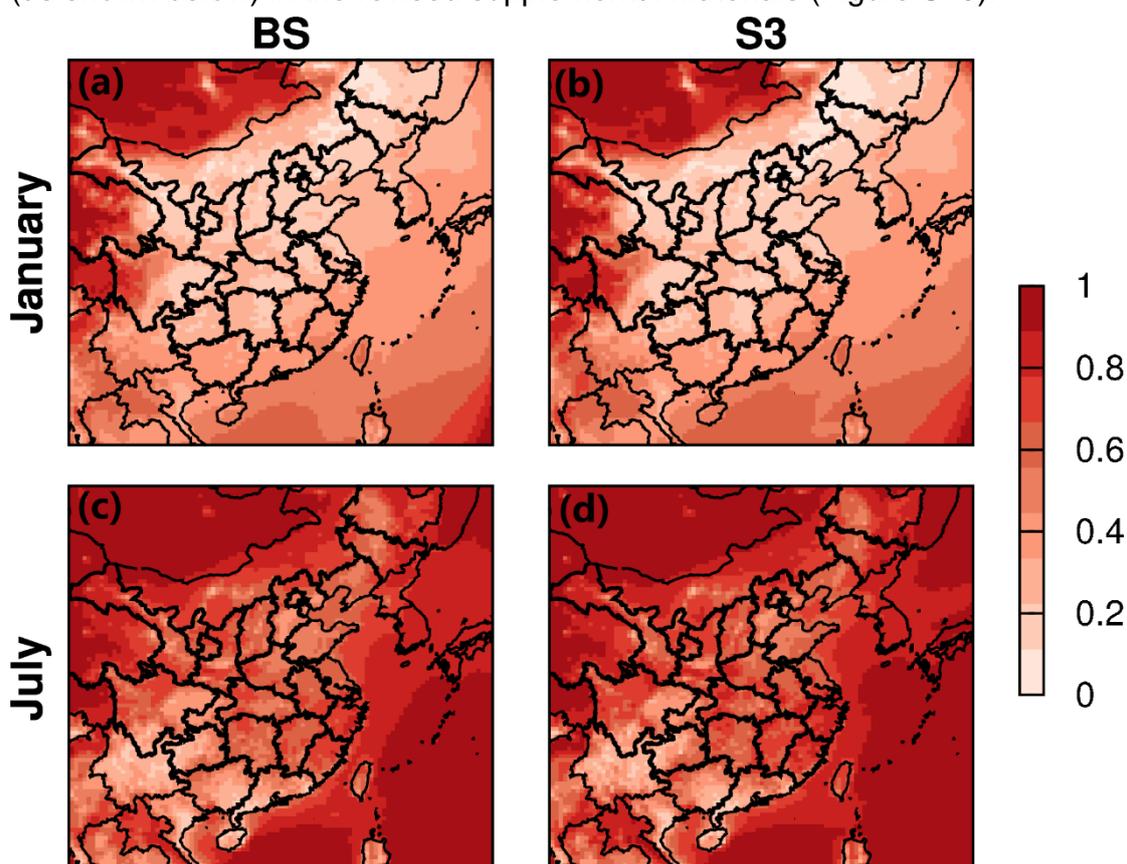


Figure R2. Averaged SOA/OA ratio from case BS and S3 during January and July of 2013.

Comment 18: Page 13 line 317: You need a zonal map to show how the water partition changes with altitude and not the total column.

Response 18: Most of the SOA and ALW_{org} retain in the lower levels of the troposphere. The information on altitude variation may not be very useful. Thus, no changes were made regarding this comment. We have also deleted this sentence in the revised manuscript to avoid confusion.

Comment 19: Page 13, lines 322-336: It is not clear how you calculate the κ_{org} in your model. This is very important for this section.

Response 19: We explained this in the revised text in L214-216:

“Since ALW_{org} in this study is calculated mechanistically using the partitioning theory, κ_{org} can be estimated by rearranging Eq3:

$$\kappa_{org} = \frac{ALW_{org}}{\rho_w V_{org}} \times \frac{1 - a_w}{a_w} \quad (\text{Eq4})$$

”

Comment 20: Figure 1: The quality of the figure is poor. It is extremely difficult to see all the plotting data and the changes due to the use of different scenarios (especially in figure 1b).

Response 20: We removed the results of BS since they are very similar to those of S3. Now the figure has been revised as follows:

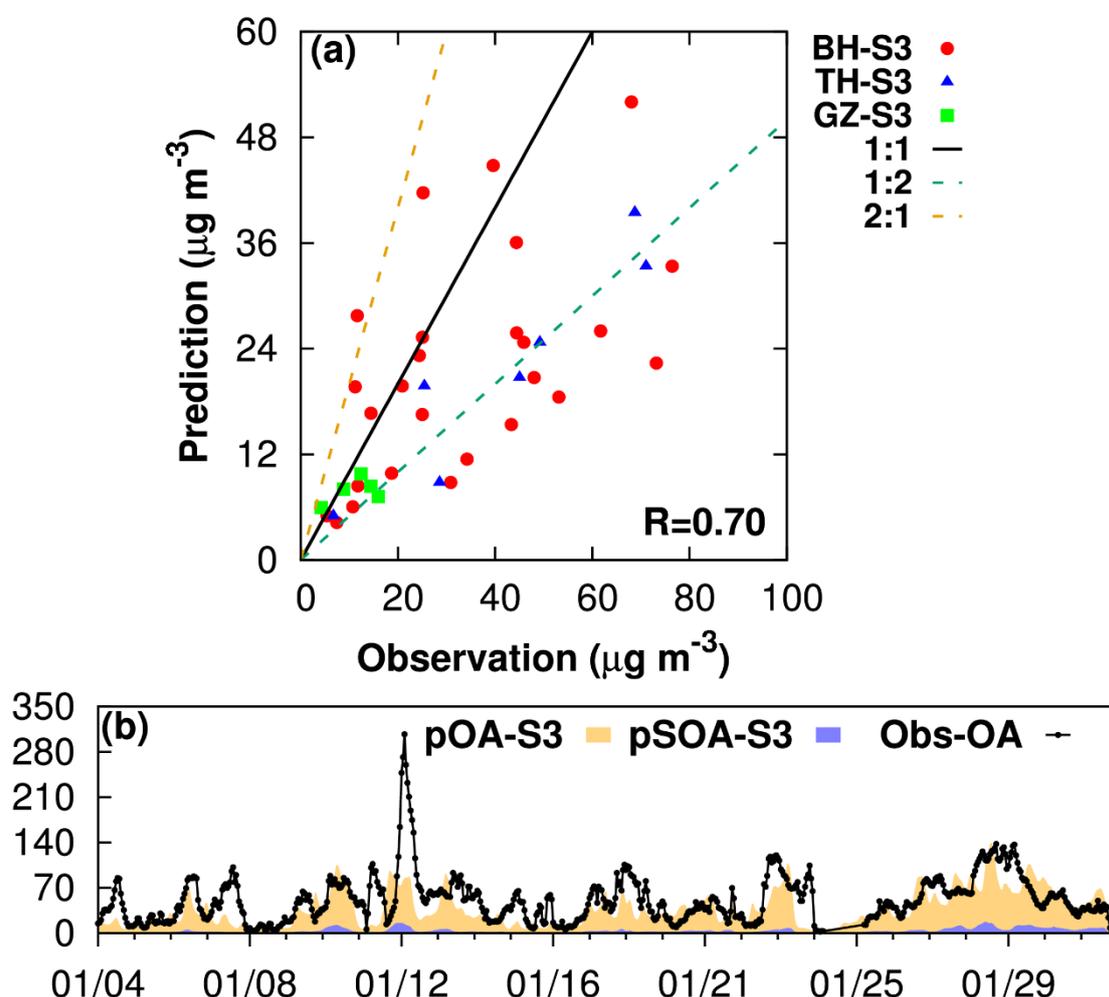


Figure 1. Comparison of (a) observed and modeled organic carbon concentration at University of Beihang (BH), Tsinghua University (TH) and Guangzhou (GZ); (b) observed organic aerosol (Obs-OA) at Beijing and predictions of total OA (pOA) and SOA (pSOA), unit is $\mu\text{g m}^{-3}$. Locations of monitoring sites are shown in Figure S1.

Comment 21: Figure 2: I found the use of daily maximum concentration in the

“difference” maps misleading. Given that you have the monthly average SOA from the basecase simulation, I would prefer to see the absolute (and relative) change of the monthly average SOA due to the use of S3 as well, and not the daily maximum. Furthermore, please add (a), (b), etc. to each subplot of the figure and add this information to the figure caption (apply this change in the rest of the figures as well).

Response 21: The monthly-averaged daily maximum differences have been replaced by the monthly-averaged differences for all the corresponding figures to reflect the general impacts on SOA and ALW_{org}. Each panel of the figure is labeled in sequence.

Comment 22: Figure 4: All the fitted correlations listed here suggest that compounds with very low or zero O:C have negative hygroscopicity. Can you comment on this limitation and include a discussion in the text?

Response 22: The relatively low values of hygroscopicity for low O:C ratio might be due to the linear regression. We also did an exponential fitting for the two variables so that the hygroscopicity falls in the range of (0,1) and is positively correlated with O:C ratio. The text and figure 4 have been revised accordingly.

In L402-407:

“In both seasons, κ_{org} approaches zero and negative values as O:C decreases, which might be due to the linear regression of κ_{org} and O:C. To avoid this, an exponential fitting of the two variables is performed so that κ_{org} falls in the range of (0,1) and is positively correlated with O:C. In this case, the fitted correlations are $\kappa_{org}=1-\exp(-(O:C/1.88)^{2.29})$ and $\kappa_{org}=1-\exp(-(O:C/1.06)^{4.50})$ for January and July of 2013, respectively.”

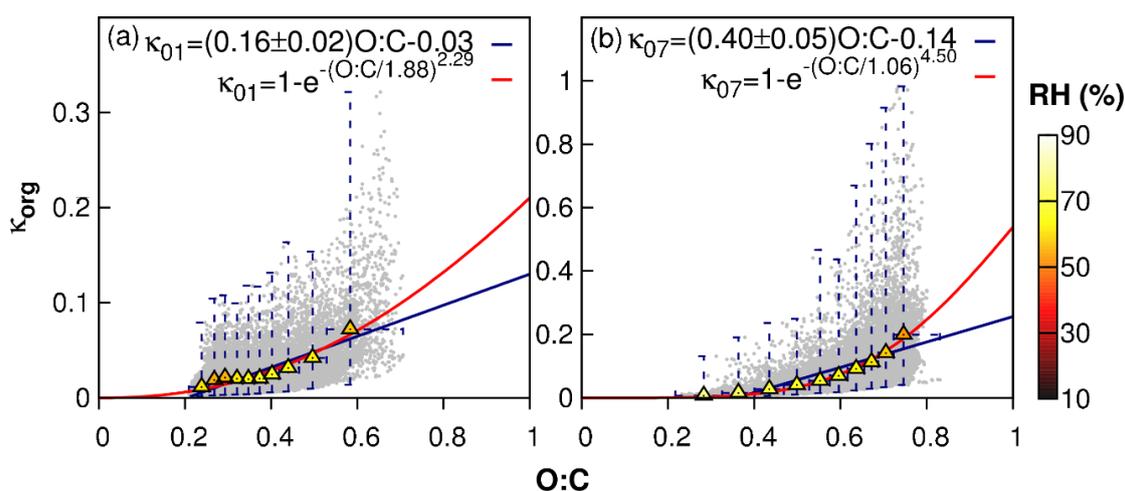


Figure 4. The correlation of hygroscopicity of organic aerosol (κ_{org}) and O:C ratio at 9 representative cities including Shenyang (SS), Beijing (BJ), Jinan (JN), Zhengzhou (ZZ), Xi’an (XA), Nanjing (NJ), Shanghai (SH), Chengdu (CD), and

Guangzhou (GZ) in January (a) and July (b) of 2013. O:C ratios are categorized into 10 bins. In each bin, the ranges of O:C and K_{org} are represented by bars. The mean values of O:C and K_{org} are represented by triangles colored by the averaged RH of each bin. The relationship between K_{org} and O:C is fitted by a linear function with reduced major axis regression (blue lines) and an exponential function (red lines), respectively. K_{01} and K_{07} represent the fitted correlation for January and July, respectively.

Comment 23: *Table S3. Please explain in the table what the fraction in the last column stands for. How have you estimated the molecular weight and fraction of the POA from unknown compounds?*

Response 23: The last column is the molar fraction of each POA surrogate. We have clarified this in the first row of this column in the table. The molecular weight and molar fraction of unknown compounds of POA have been listed in the original table already, which are 390 and 0.3, respectively.

Minor Comments

Comment 24: *The language and structure of the sentences can be substantially improved in many parts of the manuscript. Just a few examples are listed here, but I suggest revising thoroughly the wording in the whole text.*

Response 24: We thank the reviewer for pointing this out. The whole text, as well as figures and tables, have been revised carefully.

Comment 25: *Page 5, lines109-111: Please rephrase. The sentence sounds wrong.*

Response 25: The text has been revised in L109-111 as follows:

“Laboratory and field studies have observed water absorbed by SOA from a variety of precursor VOCs (Lambe et al., 2011; Zhao et al., 2016b; Asa-Awuku et al., 2010; Varutbangkul et al., 2006).”

Comment 26: *Page 5, lines117-118: Please rephrase.*

Response 26: The sentence has been revised in L116-117 as follows:

“The total water content is the summation of water associated with each solute at the same water activity.”

Comment 27: *Page 6 line 144: OC and OA abbreviations have not been used before in the main text.*

Response 27: OA abbreviation has been explained in a previous part of the revised manuscript in L120-122:

“Pye et al. (2017) found that the modeled organic aerosol (OA) improved significantly but biased high at nighttime when ALW_{org} is included in the calculation.”

Thus, this sentence has been revised in L139-141 as follows:

“The model performance was evaluated against observed meteorological parameters (temperature and relative humidity, RH) as well as PM_{2.5}, organic carbon (OC), and OA at ground monitoring sites.”

Comment 28: *Page 8, line 204: Please correct the “can estimated” to “can be estimated”.*

Response 28: The text has been revised as instructed.

Comment 29: *Page 8 line 206: I would use the word “correlate” instead of “dependent”*

Response 29: The text has been revised in L219-229 as follows:

“In many studies, κ_{org} is assumed to increase linearly with the oxidation state of OA, expressed as the O:C ratio (Massoli et al., 2010; Duplissy et al., 2011; Lambe et al., 2011).”

Comment 30: *Page 9 line 229: Change “in” with “on”*

Response 30: The text has been revised as instructed.

Comment 31: *Page 10 line 254: Change the sentence to: “In Beijing and Guangzhou, these impacts are not significant during winter”*

Response 31: The sentence has been revised in L298-300 as:

“No significant differences in OC are observed in S3 compared to BS (not shown), likely due to the biased-low SOA predicted in the current model so that limits the impact of ALW_{org} on SOA formation.”

Comment 32: *Page 11 line 274-275: Please rephrase.*

Response 32: The sentence has been revised as follows:

“The criteria of MFB and MFE followed recommendations by Boylan and Russell (2006).”

Comment 33: *Page 11 line 282: Which two areas? You have mentioned several areas in the previous sentence.*

Response 33: The sentence has been revised as follows:

“Monthly-averaged SOA concentrations in the above areas are up to 25 and 15-20 $\mu\text{g m}^{-3}$, respectively.”

Comment 34: *Page 13 line 317: “column water”. Please rephrase*

Response 34: The text has been revised in L378 as follows:

“Based on the column concentrations of ALW_{org} and ALW_{org}/SOA ratio (Figure S8),”

References

Carlton, A. G., Bhave, P. V., Napelenok, S. L., Edney, E. O., Sarwar, G., Pinder, R. W., Pouliot, G.

A., and Houyoux, M.: Model Representation of Secondary Organic Aerosol in CMAQv4.7, *Environ. Sci. Technol.*, 44, 8553-8560, 10.1021/es100636q, 2010.

Hu, J., Wang, P., Ying, Q., Zhang, H., Chen, J., Ge, X., Li, X., Jiang, J., Wang, S., Zhang, J., Zhao, Y., and Zhang, Y.: Modeling biogenic and anthropogenic secondary organic aerosol in China, *Atmos. Chem. Phys.*, 17, 77-92, 10.5194/acp-17-77-2017, 2017.

Liu, J., Shen, J., Cheng, Z., Wang, P., Ying, Q., Zhao, Q., Zhang, Y., Zhao, Y., and Fu, Q.: Source apportionment and regional transport of anthropogenic secondary organic aerosol during winter pollution periods in the Yangtze River Delta, China, *Sci. Total Environ.*, 710, 135620, <https://doi.org/10.1016/j.scitotenv.2019.135620>, 2020.

Pankow, J. F., Marks, M. C., Barsanti, K. C., Mahmud, A., Asher, W. E., Li, J., Ying, Q., Jathar, S. H., and Kleeman, M. J.: Molecular view modeling of atmospheric organic particulate matter: Incorporating molecular structure and co-condensation of water, *Atmos. Environ.*, 122, 400-408, <https://doi.org/10.1016/j.atmosenv.2015.10.001>, 2015.

Pye, H. O. T., Murphy, B. N., Xu, L., Ng, N. L., Carlton, A. G., Guo, H., Weber, R., Vasilakos, P., Appel, K. W., Budisulistiorini, S. H., Surratt, J. D., Nenes, A., Hu, W., Jimenez, J. L., Isaacman-VanWertz, G., Misztal, P. K., and Goldstein, A. H.: On the implications of aerosol liquid water and phase separation for organic aerosol mass, *Atmos. Chem. Phys.*, 17, 343-369, 10.5194/acp-17-343-2017, 2017.

1 **Impacts of water partitioning and polarity of organic compounds on**
2 **secondary organic aerosol over Eastern China**

3 Jingyi Li^{1, 2}, Haowen Zhang², Qi Ying^{3,*}, Zhijun Wu^{4, 1}, Yanli Zhang^{5,6}, Xinming
4 Wang^{5,6,7}, Xinghua Li⁸, Yele Sun⁹, Min Hu^{4, 1}, Yuanhang Zhang^{4, 1}, Jianlin Hu^{1, 2,*}

5
6 ¹ Collaborative Innovation Center of Atmospheric Environment and Equipment
7 Technology, Nanjing University of Information Science & Technology, Nanjing 210044,
8 China

9 ² Jiangsu Key Laboratory of Atmospheric Environment Monitoring and Pollution
10 Control, School of Environmental Science and Engineering, Nanjing University of
11 Information Science & Technology, Nanjing 210044, China

12 ^{3,3} [Zachry Department of Civil and Environmental Engineering](#), Texas A&M University,
13 College Station, Texas 77843-3136, USA

14 ⁴ State Key Joint Laboratory of Environmental Simulation and Pollution Control, College
15 of Environmental Sciences and Engineering, Peking University, Beijing 100871, China

16 ⁵ State Key Laboratory of Organic Geochemistry and Guangdong Key Laboratory of
17 Environmental Protection and Resources Utilization, Guangzhou Institute of
18 Geochemistry, Chinese Academy of Sciences, Guangzhou 510640, China

19 ⁶ Center for Excellence in Regional Atmospheric Environment, Institute of Urban
20 Environment, Chinese Academy of Sciences, Xiamen 361021, China

21 ⁷ University of Chinese Academy of Sciences, Beijing 100049, China

22 ⁸ School of Space & Environment, Beihang University, Beijing 100191, China

23 ⁹ State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric
24 Chemistry, Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing
25 100029, China

26 Corresponding authors:

27 Qi Ying, Email: qying@civil.tamu.edu

28 Jianlin Hu, Email: jianlinhu@nuist.edu.cn

29

30 **Abstract**

31 Secondary organic aerosol (SOA) is an important component of fine particulate matter
32 (PM_{2.5}) ~~in China~~. Most air quality models use an equilibrium partitioning method along
33 with ~~estimated~~the saturation vapor pressure (SVP) of semi-volatile organic compounds
34 (SVOCs) to predict SOA formation. However, ~~this method ignores the models typically~~
35 ~~assume that the organic particulate matter (OPM) is an ideal mixture and ignore the~~
36 partitioning of water vapor to ~~the organic aerosols and the organic phase non ideality, both~~
37 ~~of which affect the partitioning of SVOCs.~~OPM. In this study, the Community Multi-scale
38 Air Quality model (~~CMAQv5.0.1~~) ~~was used~~ CMAQ is updated to investigate the ~~above~~
39 ~~impacts~~ impacts of water vapor partitioning and non-ideality of the organic-water mixture
40 on SOA formation during winter (January) and summer (July) of 2013 over eastern China.
41 The ~~organic aerosol module was updated by incorporating water~~model treats the
42 partitioning of water vapor molecules into OPM and uses the UNIFAC model to estimate
43 the activity coefficients of species in the organic ~~particulate matter (OPM) and considering~~
44 ~~non-ideality of organic~~-water mixture. The modified model can generally capture the
45 observed surface organic carbon (OC), ~~the total organic aerosol (t)~~ with a correlation
46 coefficient R of 0.7, and the surface OA) ~~and diurnal variation of PM_{2.5} at ground sites.~~
47 with the mean fractional bias (MFB) and mean fractional error (MFE) of -0.28 and 0.54,
48 respectively. SOA concentration shows significant seasonal and spatial variations, with
49 high ~~concentration levels~~concentrations in the North China Plain (NCP), Central China and
50 Sichuan basin (SCB) ~~areas~~regions during winter (up to 25 µg m⁻³) and in the Yangtze River
51 Delta (YRD) during summer (up to ~~1216~~ µg m⁻³). ~~When water partitioning is included in~~
52 In winter, SOA ~~concentrations increase~~ decreases slightly in the updated model, with the
53 monthly-averaged ~~daily maximum~~ relative ~~differenee~~change of 10-20% in the highly
54 concentrated areas, mainly due to organic-water interactions. The monthly-averaged
55 concentration of SOA increases greatly in summer, by 20-50% at the surface and ~~10-30%~~
56 ~~for 60%~~ in the whole column, ~~mostly~~. The increase of SOA is mainly due to the increase

57 in biogenic SOA in inland areas and anthropogenic SOA. ~~The increase in SOA is more~~
58 ~~significant in summer, by 20–90% at the surface and 30–70% for the whole column. The~~
59 ~~increase of SOA over the land is mostly due to biogenic SOA while the increase of SOA~~
60 ~~over the in coastal regions is related with that of anthropogenic origin. Further analysis~~
61 ~~of two representative cities, Jinan and Nanjing, shows that changes of SOA are favored~~
62 ~~under hot and humid conditions. The increases in SOA cause a 12% elevation in the areas.~~
63 As a result, the averaged aerosol optical depth (AOD) is increased by up to 10% and 15%
64 enhancement in the cooling effect of aerosol radiative forcing (ARF) is enhanced by
65 up to 15% over YRD in summer. The aerosol liquid water content associated with OPM
66 (ALW_{org}) at the surface is relatively high ~~over the land in inland areas~~ in winter and over
67 the ocean in summer, with the monthly-averaged ~~daily maximum concentration~~ of 2–90.5–
68 3.0 and 5–127 $\mu\text{g m}^{-3}$, respectively. ~~By using the~~ The hygroscopicity parameter (κ) of OA
69 based on the κ -Köhler theory, we calculated the hygroscopicity of OA with is
70 determined using the modeled ALW_{org}, ~~finding that the~~. The correlation of κ with O₃: C
71 ratio varies significantly across different cities and seasons. Water Analysis of two
72 representative cities, Jinan (in NCP) and Nanjing (in YRD), shows that the impacts of water
73 partitioning and non-ideality of the organic-water mixture on SOA are sensitive to
74 temperature, relative humidity (RH), and the SVP of SVOCs. The two processes exhibit
75 opposite impacts on SOA in eastern China. Water uptake increases SOA by up to 80% in
76 the organic phase, while including non-unity activity coefficients decreases SOA by up to
77 50%. Our results indicate that both water partitioning into OPM only promotes SOA
78 formation, while non-ideality of organic-water mixture only leads to decreases in SOA in
79 most regions of eastern China. Water partitioning into OPM and the activity coefficients of
80 the condensed organics should be considered in ~~air quality models in~~ simulating SOA
81 formation from gas-particle partitioning, especially in hot and humid environments.

82
83 **Keywords:** SOA, non-ideality, water partitioning, hygroscopicity

84

85 **1 Introduction**

86 Secondary organic aerosol (SOA) is formed via a complex interaction of volatile organic
87 compounds (VOCs) with oxidants and primary particles emitted from anthropogenic and
88 biogenic sources in the atmosphere. As an important component of fine particulate matter
89 (PM_{2.5}), SOA can cause severe air pollution in urban and suburban areas (Huang et al.,
90 2014) and exhibit adverse health effects (~~Polichetti et al., 2009; Feng et al., 2016; Xing et~~
91 ~~al., 2016; Atkinson et al., 2014).~~(Atkinson et al., 2014). SOA also plays an important role
92 in new particle formation and particle growth (~~Man et al., 2015; Zhang et al.,~~
93 ~~2011; Wiedensohler et al., 2009; Yue et al., 2011; Liu et al., 2014; Ehn et al., 2014; Huang et~~
94 ~~al., 2019; Jokinen et al., 2015) and further contribute that further contribute~~ to the
95 enhancement of cloud condensation nuclei (CCN) (~~Yue et al., 2011; Wiedensohler et al.,~~
96 ~~2009; Liu et al., 2014; Jokinen et al., 2015).~~(Wiedensohler et al., 2009; Ehn et al., 2014).
97 This will, in turn, impact the atmospheric aerosol burden, precipitation and water
98 circulation, solar radiation budget, and climate (~~Rosenfeld et al., 2008; Spracklen et al.,~~
99 ~~2011; Quaas et al., 2008; Ramanathan et al., 2001; Hatzianastassiou et al., 2007; Hegerl et al.,~~
100 ~~2015). However, the mechanisms of these~~(Ramanathan et al., 2001). However, the extents
101 ~~of those~~ influences are not well understood so far, due to the high uncertainties associated
102 with the formation and physical and chemical properties of SOA (~~Shrivastava et al.,~~
103 ~~2017).~~(Shrivastava et al., 2017). Large gaps still exist in SOA mass ~~loadings~~loading and
104 properties between ~~model—estimates~~models and ~~laboratory—and—field~~
105 ~~measurements~~observations (Gentner et al., 2017; Ervens et al., 2011; Hayes et al.,
106 2015)(Gentner et al., 2017; Ervens et al., 2011; Hayes et al., 2015). Therefore, it is crucial
107 to explore and resolve this issue to improve our knowledge of the roles of SOA in the
108 environment, human health, and climate.

109 Gas-particle partitioning of semi-volatile and low-volatile organic compounds
110 (SVOCs and LVOCs) generated from VOC oxidation is an important pathway of SOA

111 formation. In most current chemical transport models (CTMs), this process is treated as an
112 equilibrium partitioning process that depends on the mass concentration of the organic
113 particulate matter (OPM), ambient temperature, (T), the mean molecular weight of OPM,
114 and the volatility of ~~pure~~ condensed organics ~~(Pankow, 1994)~~(Pankow, 1994). The
115 ~~volatilities~~formation of condensed organic products ~~from a certain precursor VOC are~~
116 ~~either is commonly~~ represented by ~~that of several lumped surrogates based on surrogate~~
117 SVOCs in a 2-product model with volatilities and SVOC yields fitted to chamber
118 ~~experiments (2-product model) (Odum et al., 1996)~~(Odum et al., 1996) ~~or fitted into~~
119 ~~different bins of a fixed volatility range (usually 0.01–10⁵ μg m⁻³) (volatility basis set model,~~
120 ~~VBS model) (Donahue et al., 2006). Although the above.~~ To better represent the volatility
121 of primary organic aerosol (POA) and the multi-generation oxidation of SVOCs to a wider
122 range, Donahue et al. (2006) proposed the volatility basis set (VBS) model in which the
123 mass yields of SVOCs are fitted to a fixed number of volatility bins (usually 0.01–10⁵ μg
124 m⁻³). The VBS model has been adopted by several CTMs (such as WRF-Chem, GEOS-
125 Chem, etc.).

126 Although the models can capture the general trend of SOA evolution and mass
127 concentration to some extent ~~(Slowik et al., 2010; Li et al., 2017a; Baek et al.,~~
128 ~~2011; Bergström et al., 2012; Woody et al., 2016; Heald et al., 2006)~~(Li et al., 2017a;
129 Bergström et al., 2012; Woody et al., 2016), ~~both of them neglected,~~ two key factors
130 ~~that currently neglected in models~~ may lead to biases: 1) the molecular structures and
131 interactions of functional groups (-OH, -C=O, -COOH, etc.) of condensed organics (non-
132 ideality); 2) partitioning of water vapor, ~~an~~the most abundant atmospheric constituent
133 besides O₂ and N₂, to OPM. The non-ideality alters the volatility of condensed organics,
134 and thus their contributions to the total SOA mass loading (Cappa et al., 2008). Water
135 partitioning into OPM can reduce the partial pressure of organics ~~and lead~~due to ~~increase~~
136 ~~in SOA mass, which is called the~~ Raoult's Law effect ~~(Prisle et al., 2010)~~(Prisle et al.,
137 2010). ~~This impact~~ and lead to increases in SOA mass. The amount of aerosol liquid water

138 associated with organics (ALW_{org}) may vary for different SOA-precursors (~~Healy et al.,~~
139 ~~2009; Prisle et al., 2010~~)(Healy et al., 2009; Prisle et al., 2010). The above two aspects will
140 not only affect the chemical composition of SOA but also the inorganic portion (~~Ansari~~
141 ~~and Pandis, 2000; Meyer et al., 2009~~) and optical properties (~~Liu and Wang, 2010; Denjean~~
142 ~~et al., 2015~~)(Ansari and Pandis, 2000) and optical properties (Denjean et al., 2015) of
143 aerosols.

144 Laboratory and field studies have ~~confirmed the fact that~~observed water absorbed by
145 SOA (~~quantified as hygroscopicity, κ~~) from a variety of precursor VOCs (~~Lambe et al.,~~
146 ~~2011; Zhao et al., 2016b~~)(Lambe et al., 2011; Zhao et al., 2016b; Asa-Awuku et al.,
147 2010; Varutbangkul et al., 2006; Varutbangkul et al., 2006). The hygroscopicity of SOA ~~is~~
148 ~~highly, quantitatively described by the hygroscopicity parameter, κ , is~~ correlated with the
149 oxygen-to-carbon ratio (O:C) and increases with more oxidized SOA during photochemical
150 aging (~~Poulain et al., 2010; Wang et al., 2014; Lambe et al., 2011; Tritscher et al.,~~
151 ~~2011a; Zhao et al., 2016b; Massoli et al., 2010; Tritscher et al., 2011b; Duplissy et al.,~~
152 ~~2014~~)(Lambe et al., 2011; Zhao et al., 2016b). The OPM-associated water partitioning can
153 be estimated using the κ -Köhler theory under the Zdanovskii-Stokes-Robinson (ZSR)
154 assumption of no interactions between any constituents in aerosols (~~Petters and Kreidenweis,~~
155 ~~2007~~)(Petters and Kreidenweis, 2007). The total water content is the
156 ~~summarization~~summation of water associated with each ~~constitutes~~solute at the same
157 ~~RH-water activity~~. Guo et al. (~~2015~~)(2015) found that this simplified method, along with
158 the ISORROPIA model which is used to predict aerosol liquid water (ALW) associated
159 with the inorganic portion of aerosols, ~~reproduced~~could reproduce the observed total ALW
160 in the ambient environment. Pye et al. (~~2017~~)(2017) ~~applied this approach along with a~~
161 ~~parameterization of overall κ based on O:C ratio and a simplified method to estimate~~
162 ~~activity coefficients of organics and~~ found that the modeled organic aerosol (OA and
163 ALW are) improved during daytime significantly but ~~still~~-biased lowhigh at nighttime-
164 ~~Shortcomings still exist in the above method for water associated organics (when ALW_{org})~~

165 is included in the calculation. However, as ~~interactions between the interaction among~~
166 ~~organic species~~ and between organics and water in the organic-water mixture ~~are not~~
167 ~~considered, which~~ has been shown to play an important role in SOA formation and water
168 partitioning to OPM (~~Kim et al., 2019~~)(~~Kim et al., 2019~~). ~~A representation of water~~
169 ~~partitioning along with SVOCs with consideration of water-organic and organic-organic~~
170 ~~interactions in CTMs showed significant influences in SOA and ALW in the eastern U.S.~~
171 ~~where biogenic SOA dominated in OA and the internal mixing assumed for the aerosol~~
172 (~~Pankow et al., 2015; Jathar et al., 2016~~).

173 ~~China has been suffering from severe PM_{2.5} pollution especially in the eastern region~~
174 ~~with fast urbanization and economic development,~~ the ALW_{org} estimated by the κ-Köhler
175 theory and its impact on SOA might not be accurate. Using UNiversal Functional Activity
176 Coefficient (UNIFAC) method (Fredenslund et al., 1975) for calculating activity
177 coefficients of the organic-water mixture, it was found that in the eastern U.S., where
178 biogenic SOA dominated the OA, considering ALW_{org} leads to a significant increase in
179 predicted SOA (Guo et al., 2014; Fu and Chen, 2017; Yang et al., 2016)(Pankow et al., 2015;
180 Jathar et al., 2016). ~~The secondary portion has been proved to be dominated in PM_{2.5} and~~
181 ~~organic aerosol increases during haze events (Huang et al., 2014; Sun et al., 2019). In~~
182 ~~addition,~~.

183 China has been suffering from severe PM_{2.5} pollution especially in the eastern region
184 with fast urbanization and economic development (Fu and Chen, 2017). SOA is a very
185 important component of PM_{2.5} in China that contributes about 20-50% (Li et al., 2017b).
186 The fraction of SOA in OA increases during haze events (Huang et al., 2014; Sun et al.,
187 2019). Previous modeling studies ~~indicated in China indicate~~ that SOA was underpredicted
188 ~~in this region (Wang et al., 2011; Lin et al., 2016; Jiang et al., 2012)(Lin et al., 2016; Jiang~~
189 ~~et al., 2012)~~ and the impacts of the non-ideality and water-OPM partitioning on modeled
190 SOA have not been evaluated.

191 In this study, regional simulations of SOA during January and July of 2013 over
192 eastern China ~~under several scenarios~~ were conducted to investigate the seasonal variation
193 of SOA due to water partitioning into OPM. ~~Model performances were firstly~~The model
194 performance was evaluated against observed meteorological parameters (temperature and
195 relative humidity, RH) as well as PM_{2.5}, organic carbon (OC₂), and OA at ground
196 monitoring sites. ~~Then, the~~The regional and seasonal impacts on SOA, ALW_{org}, and ~~water~~
197 ~~content~~properties of aerosols were quantified. ~~Factors related to the impacts on SOA,~~Lastly,
198 influences of the results by several factors including ~~sources of precursors, chemical~~
199 ~~compositions and meteorological conditions were further parameters, estimations of~~
200 saturation vapor pressures (SVP) of condensed organics, and the individual impacts of
201 ALW_{org} and non-ideality of the organic-water mixture on SOA prediction were analyzed.
202 ~~Lastly, the impacts on aerosol optical properties and hygroscopicity were investigated.~~

203 **2 Methodology**

204 **2.1 Model description**

205 The Community Multi-scale Air Quality model (CMAQ v5.0.1) coupled with a modified
206 SAPRC-11 was used in this study. Model configurations were largely based on that used
207 by Hu et al. (2016) as summarized below. Firstly, SAPRC-11 was expanded for a more
208 detailed treatment of isoprene oxidation and tracking dicarbonyl products (glyoxal and
209 methylglyoxal) ~~products~~ from different groups of major precursors (~~Ying et al., 2015~~);
210 ~~Secondly, heterogeneous formation of secondary nitrate and sulfate from NO₂ and SO₂~~
211 ~~reaction on particle surface (Ying et al., 2014), and (Ying et al., 2015).~~ Secondly, SOA from
212 isoprene epoxydiols (IEPOX), methacrylic acid epoxide (MAE) and dicarbonyls through
213 surface-controlled irreversible reactive uptake were added (~~Li et al., 2015; Pankow et al.,~~
214 ~~2015~~)(Hu et al., 2017; Li et al., 2015; Liu et al., 2020; Ying et al., 2015) ~~were added;~~
215 ~~Thirdly, Thirdly, the heterogeneous formation of secondary nitrate and sulfate from NO₂~~
216 ~~and SO₂ reaction on the particle surfaces (Ying et al., 2014) were added, which is an~~
217 important source of secondary inorganic aerosols (Zheng et al., 2015) and improves model

218 estimates of nitrate and sulfate (Qiao et al., 2018; Shi et al., 2017). Fourthly, SOA yields
219 were corrected for vapor wall loss (Zhang et al., 2014)(Zhang et al., 2014). Impacts of the
220 above updates on model performances have been extensively discussed in the cited work
221 and will not be further investigated in the current study.

222 ~~Two types of SOA were considered in the current model, “semi-volatile” (SV) portion~~
223 ~~that formed via equilibrium absorption-partitioning of SVOCs, and “non-volatile” (NV)~~
224 ~~portion that formed via direct oxidation of aromatics at low NO_x, isoprene oxidation under~~
225 ~~acidic conditions, reactive uptake of dicarbonyls, IEPOX and MAE, and oligomers. The~~
226 ~~SV SOA module mostly based on that of Pankow et al. The SOA module mostly follows~~
227 ~~Pankow et al. (2015)(2015) with several updates in the treatment of primary organic aerosol~~
228 ~~(POA) by including it in the non-ideality calculation of the organic-water mixture. The~~
229 ~~mass distribution of SVOCs between the gas-phase and particle-phase follows the equation:~~

230 . Two types of SOA as traditionally treated in CMAQ were considered, “semi-volatile”
231 (SV) portion that formed via equilibrium absorption-partitioning of SVOCs, and “non-
232 volatile” (NV) portion that includes the oligomers and SOA formed via direct oxidation of
233 aromatics at low-NO_x. SOA from dicarbonyls, IEPOX, and MAE were formed by
234 irreversible reactive uptake and categorized as NV-SOA in the current model as well. Some
235 studies investigated SOA from glyoxal, methylglyoxal, and IEPOX using detailed
236 reactions and reversible pathways in models or observed as reversible processes in chamber
237 experiments, leading to a relatively lower SOA yield compared to the surface-controlled
238 irreversible uptake (Lim et al., 2013; Knote et al., 2014; Galloway et al., 2009; El-Sayed et
239 al., 2018; Budisulistiorini et al., 2017). The non-volatile assumption used in this paper
240 allows an upper-limit estimation of the importance of these additional SOA formation
241 pathways. POA was treated as non-volatile and non-reactive. The mass distribution of
242 SVOCs between the gas-phase and particle-phase follows the equation:

$$K_{p,i} = \frac{F_i}{M \cdot A_i} \quad (\text{Eq-1})$$

243 where $K_{p,i}$ ($m^3 \mu g^{-1}$) is the gas-particle partitioning constant for compound i ,
 244 F_i ($\mu g m^{-3}$) is the concentration of species i in the particle-phase, A_i ($\mu g m^{-3}$) is the
 245 concentration of species i in the gas-phase, and M ($\mu g m^{-3}$) is the total mass
 246 concentration of the absorbing organic phase (i.e. OPM). The gas-particle partitioning
 247 constant $K_{p,i}$ is dependent on the chemical composition of the absorbing organic
 248 phase-OPM. Pankow et al. (1994)(1994) derived $K_{p,i}$ for SVOCs partitioning into an
 249 absorbing organic phase as:

$$K_{p,i} = \frac{RT}{10^6 \overline{MW} \xi_i p_{L,i}^o} \quad (\text{Eq-} \quad \text{2Eq2})$$

250 where $p_{L,i}^o$ (atm) is the saturation vapor pressure SVP of the pure compound i at temperature
 251 T (K), ξ_i is the activity coefficient of species i in the absorbing organic phase, \overline{MW} (g mol⁻¹)
 252 ¹) is the average molecular weight of the absorbing organic phase OPM, R (8.314 J mol⁻¹
 253 K⁻¹) is the gas constant, and 10^6 is used to convert the unit units to $m^3 \mu g^{-1}$.

254 There are 12 lumped SVOCs generated by oxidation of alkanes, alkenes, and
 255 aromatics oxidized under different NO_x conditions (and 8 NV organic products as listed in
 256 Table S1), and Table S2. More details about the lumped precursors such as formation
 257 conditions (“high” and “low” NO_x), lumping species and method, and yields from parent
 258 VOCs can be found in Carlton et al. (2010) and summarized in SI. Activity coefficients of
 259 SVOCs were calculated based on the composition of absorbing organic phase OPM using
 260 the UNiversal Functional Activity Coefficient (UNIFAC) method (Fredenslund et al.,
 261 1975), method, with assigned carbon number (n_c), functional groups and energy interaction
 262 parameters to both SV and NV compounds (Pankow et al., 2015)(Pankow et al., 2015).
 263 The UNIFAC model is one of the commonly used models that activity coefficients of
 264 condensed organics and their interactions with water can be estimated. This method has
 265 been adopted to investigate the impacts of non-ideality and water-OPM partitioning into
 266 OPM on SOA for different precursors in box models (Seinfeld et al., 2001; Bowman and
 267 Melton, 2004)(Seinfeld et al., 2001; Bowman and Melton, 2004) and CTMs (Jathar et al.,

268 ~~2016;Pankow et al., 2015;Kim et al., 2019)~~(Pankow et al., 2015; Kim et al., 2019). The
269 ~~primary organic aerosols (In the current model, POA)~~ was assumed to have a bulk
270 composition of ten categories of surrogate species (Table S3), as used by Li et al.
271 ~~(2015)(2015)~~. POA is also involved in the calculation of activity coefficients for the
272 ~~organics in the condensed phase.organic-water mixture~~. Detailed information about the
273 surrogate species including ~~the~~their structures and properties can be found in Li et al.
274 ~~(2015)(2015)~~ and references therein.

275 In addition to organic compounds, water partitioning into OPM is enabled according
276 to Eq 1 and Eq 2. In such a case, the absorbing phase in Eq 1 includes both organic aerosols
277 and water ~~partitioning into~~associated with OPM. As water ~~considered~~condenses in the
278 absorbing organic phase, it will further alter the molar fraction of each composition, the
279 activity coefficient of SVOCs and the SV-SOA mass concentrations as a result. In the
280 current model, we assumed no interactions between the inorganic and organic phases.

281 ~~As the water partitioning into OPM is highly correlated with the hygroscopicity of~~
282 ~~aerosols (κ), their correlation can be expressed by the κ -Köhler theory with Kelvin effect~~
283 ~~neglected (Peter et al., 2006):~~

284 2.2 Estimating the hygroscopicity of OA

285 Based on the κ -Köhler theory with linearly additive hygroscopic behavior of each
286 component of the mixed particle, ALW_{org} is related to the hygroscopicity parameter for the
287 organic mixture (κ_{org}) by Eq3 (Petters and Kreidenweis, 2007):

$$ALW_{org} = V_{org} = \rho_w V_{org} \kappa_{org} \frac{a_w}{1 - a_w} \quad (\text{Eq3})$$

288 where ρ_w is the density of water (assumed to be 1 g cm^{-1}), V_{org} is the volume concentration
289 of ~~organic~~organics, and a_w is the water activity (assumed to be the same as RH).

290 ~~Take~~Since ALW_{org}

291 in this study is calculated mechanistically using the partitioning theory, κ_{org} can be
292 estimated by rearranging Eq3:

$$\kappa_{org} = \frac{ALW_{org}}{\rho_w V_{org}} \times \frac{1 - a_w}{a_w} \quad (\text{Eq4})$$

293 V_{org} can be estimated from the modeled mass concentration of OA, assuming the density
 294 of organic aerosol OA to be 1.2 g cm⁻³ (Li et al., 2019), the hygroscopicity of the total OA
 295 can be estimated. This simplified method can be used to estimate OPM associated water.

296 In many studies, κ_{org} is assumed to increase linearly with the oxidation state of OA,
 297 expressed as the O:C ratio (Guo et al., 2015; Li et al., 2019) (Massoli et al., 2010; Duplissy
 298 et al., 2011; Lambe et al., 2011). In addition, the hygroscopicity of organic aerosol is
 299 dependent on the degree of oxygenation, showing a positive linear relationship with the
 300 O:C ratio (Massoli et al., 2010; Duplissy et al., 2011; Lambe et al., 2011; Hong et al., 2018; Li
 301 et al., 2019). We therefore estimated the correlation of κ and O:C ratio at 9 representative
 302 cities during January and July. The correlation of κ_{org} and O:C ratio at 9 representative
 303 cities was evaluated during January and July of 2013, with the reduced major axis
 304 regression method (Ayers, 2001). The O:C ratio of the total OA was calculated as
 305 following using Eq5:

$$O:C = \sum_{i=1}^n f_i (O:C)_i \quad (\text{Eq5})$$

306 where f_i and $(O:C)_i$ are the molar fraction and O:C ratio of organic aerosol component
 307 i. For POA, a fixed molar fraction and composition has been were assumed following Li et
 308 al. (2015) (2015). For SOA, the O:C ratio was estimated by their OM:OC ratio (Simon and
 309 Bhave, 2012):

310 . For SOA, the O:C ratio was calculated by using their organic matter to organic carbon
 311 ratio (OM:OC) following Simon and Bhave (2012):

$$O:C = \frac{12}{15} \equiv \frac{12}{15} (OM:OC) - \frac{14}{15} \quad (\text{Eq6})$$

312 OM:OC ratio of each SOA component follows Pye et al. (2017).

313 The OM:OC ratio of each SOA component follows Pankow et al. (2015) as shown in Table
 314 S1-S2.

315 2.3 Model application

316 The simulation domain has a horizontal resolution of 36 km × 36 km (100 × 100 grids)
317 and a vertical structure of 18 layers up to 21 km, which covers eastern China as shown in
318 Figure S1. Anthropogenic emissions were generated from the Multi-resolution Emission
319 Inventory for China (MEIC) (~~Zhang et al., 2009; Li et al., 2014; Zheng et al., 2014; Liu et~~
320 ~~al., 2015~~)(Zhang et al., 2009) v1.0 with a 0.25° × 0.25° resolution
321 (<http://www.meicmodel.org>) for China, and the Regional Emission inventory in Asia
322 version 2 (REAS2) (Kurokawa et al., 2013) with a 0.25° × 0.25° resolution
323 (<http://www.nies.go.jp/REAS/>) for the rest of the domain. Biogenic emissions were
324 generated by the Model for Emissions of Gases and Aerosols from Nature (MEGAN) v2.1,
325 with the leaf area index (LAI) from the 8-day Moderate Resolution Imaging
326 Spectroradiometer (MODIS) LAI product (MOD15A2) and the plant function types (PFTs)
327 from the Global Community Land Model (CLM 3.0). Open biomass burning emissions
328 were generated from the Fire INvnetory from NCAR (FINN) (~~Wiedinmyer et al.,~~
329 ~~2011~~)(Wiedinmyer et al., 2011). Dust and sea salt emissions were generated ~~in line~~online
330 during CMAQ simulations. The total emissions of major SOA precursors and their spatial
331 distributions are shown in Table S4 and Figure S2. Meteorological fields were generated
332 using the Weather Research and Forecasting (WRF) model v3.6.1 with initial and boundary
333 conditions from the NCEP FNL Operational Model Global Tropospheric Analyses dataset.
334 More details about the model application can be found in Hu et al. (2016).

335 Four scenarios are investigated in this study. The base case (BS) ~~that applied~~applies
336 the default secondary organic aerosol module of CMAQ; ~~the water v5.0.1. In this case (S1)~~
337 ~~that only, no~~ water partitioning into OPM ~~was is~~ considered; ~~the UNIFAC case (S2) that~~
338 ~~effects. Lumped semi-volatile products from the oxidation of molecular structure of the~~
339 ~~primary and secondary various precursors partition into a single organic species were~~
340 ~~included;~~phase, which is considered as an ideal mixture of POA and SOA with $\gamma_{org}=1$.
341 The water case (S1) includes water partitioning into OPM, which is again considered as an

342 ideal solution ($\gamma_{org}=1$ and $\gamma_{H_2O} = 1$). The UNIFAC case (S2) considers the interaction
343 between organic constituents with UNIFAC calculated activity coefficients ($\gamma_{org} \neq 1$) but
344 does not allow water partitioning into OPM. The combined case (S3) ~~that~~ allows both water
345 partitioning and interactions between all constituents (including water and organics) using
346 UNIFAC calculated activity coefficients ($\gamma_{org} \neq 1$ and $\gamma_{H_2O} \neq 1$). The results of BS and
347 S3 are used to examine the overall impacts of water partitioning into OPM and polarity of
348 organics on SOA and ALW_{org} , as shown in Section 3.1-3.4. The separate influences of
349 those two processes on SOA from S1 and S2 ~~were combined together.~~ are discussed in
350 Section 3.5.

351 **3 Results and discussion**

352 **3.1 Model evaluation**

353 ~~Temperature~~ The meteorological inputs and relative humidity (emissions have been used in
354 several previous publications. Model performance on meteorological parameters
355 (temperature and RH), gaseous species and gas and aerosol concentrations have been
356 extensively evaluated (Hu et al., 2016; Hu et al., 2017; Qiao et al., 2018; Shi et al., 2017).
357 A summary of the model performance related to this study is provided below. Observed
358 meteorological data were obtained from the National Climatic Data Center
359 (<ftp://ftp.ncdc.noaa.gov/pub/data/noaa>). Observations of OC at two urban locations,
360 Beijing (Cao et al., 2014; Wang et al., 2015) and Guangzhou (Lai et al., 2016) and OA in
361 Beijing (Sun et al., 2014) during January of 2013 as well as surface $PM_{2.5}$ at several
362 monitoring sites during July of 2013 from China National Environmental Monitoring
363 Center (<http://113.108.142.147:20035/emcpublish/>) were used to evaluate model estimates
364 of aerosols. Details of measurement methodology and uncertainties of observations are
365 listed in the corresponding references.

366 Temperature and RH are the two meteorological factors that affect SOA formation.
367 Table 1 ~~shows the comparison of~~ lists model statistics of mean observation (OBS), mean
368 prediction (PRE), mean bias (MB), gross error (GE) and correlation coefficient (R) based

369 on WRF predictions and observations at monitoring sites located in 8 sub-regions of the
370 domain (Figure S1). ~~Observed data are accessible from the National Climatic Data Center~~
371 ~~at <ftp://ftp.ncdc.noaa.gov/pub/data/noaa>. Temperature and RH are~~ during January and
372 July of 2013. The benchmarks for the MM5 model (another meteorology model) of 4-12km
373 horizontal resolution suggested by Emery et al. (2001) are also listed in the table. Details
374 of monitoring sites in the 8 sub-regions are listed in Table S5. Overall, WRF tends to
375 underestimate both temperature and RH. The model shows better agreement with observed
376 temperature as R is higher than that of RH. Both temperature and RH are well captured by
377 ~~WRF~~ the model in YRD, the Pearl River Delta (PRD), and the central regions of China (the
378 major regions of eastern China). In these regions, MB and GE of temperature are -1.2~0.7
379 K and 1.8~2.6 K, respectively, which are -11.8~5.6% and 9.2~16.8% for RH, respectively.
380 Model estimates of daily organic carbon (OC) from ~~the BS case~~ were S3 are compared with
381 measurements at monitoring sites in Beijing and Guangzhou ~~during the winter~~ in January
382 of 2013 (Figure 4(a)). Overall, 1a). The factors used to convert SOA to OC (SOC) are listed
383 in Table S1-S2. OC from POA (POC) is directly predicted by the model. Generally, the
384 ratio between modeled and observed OC concentration falls in the range of 1:2 to 2:1, with
385 ~~a correlation coefficient R of 0.70. The model tends to underestimate OC, especially in~~
386 ~~Beijing on highly polluted days (by 37~48%). No significant improvements to modeled~~
387 ~~OC were observed in S3. The impacts of water co-condensation and polarity of organic~~
388 ~~condensed species on SOA exhibit strong seasonal and spatial features, which are further~~
389 ~~discussed in Section 4. The impacts in Beijing and Guangzhou are not significant during~~
390 ~~winter.~~ an R-value of 0.7. The model tends to underestimate OC on high concentration
391 days. Overall, the mean fractional bias (MFB) and mean fractional error (MFE) of OC are
392 -0.20 and 0.27, within the criteria (MFB ≤ ±0.6; MFE ≤ 0.75) suggested by EPA (2007).
393 The bias in OC might be due to ~~under-estimated~~ underestimated POA emissions and ~~under-~~
394 ~~predicted~~ underpredicted SOA in CMAQ from missing precursors (Hu et al., 2017; ~~Zhao et~~
395 ~~al., 2016a); Zhao et al., 2016a).~~—

396 ~~The model estimate of OA was further investigated. As~~. No significant differences in
397 OC are observed in S3 compared to BS (not shown in), likely due to the biased-low SOA
398 predicted in the current model that limits the impact of ALW_{org} on SOA formation.

399 The underestimate of SOA can be seen from Figure 4(b),1b as well. CMAQ can well
400 capture the observed diurnal variation of OA ~~at~~in Beijing during wintertime, except for the
401 underestimates of peak values. ~~A better agreement~~ The correlation coefficient of modeled
402 to observed OA is 0.55. We find a 25% underestimate of OA on average. Better agreement
403 between the model and the observations is ~~observed~~shown on non-polluted days (daily-
404 averaged concentration less than 75 $\mu\text{g m}^{-3}$). ~~The monthly averaged mean fractional bias~~
405 ~~(MFB) and mean fractional error (MFE) are -0.13 and 0.27, respectively. POA is~~ The mean
406 fractional bias (MFB) and mean fractional error (MFE) of polluted days are -0.38 and 0.64,
407 respectively, which are worse than that of the non-polluted days (-0.26 for MFB and 0.52
408 for MFE). The overall MFB and MFE of OA during January are -0.28 and 0.54, within the
409 criteria (MFB \leq \pm 0.6; MFE \leq 0.75) suggested by EPA (2007). Again, no apparent changes
410 of SOA nor OA are observed between case S3 and BS (not shown), since POA is predicted
411 to be the primary contributor to OA at Beijing in winter, accounting for 88% due to aging
412 of POA not treated in the current model. The fraction of SOA is small, resulting in little
413 impacts on SOA by in the current model, with an averaged SOA/POA ratio of 0.12. This
414 ratio is much lower than the field observation of about 0.45-1.94 (Zhao et al., 2019; Sun et
415 al., 2013; Sun et al., 2016). The bias might be due to the missing SOA converted by
416 partitioning and aging of semi-volatile POA as well as oxidation from intermediate volatile
417 organic compounds (IVOCs) and VOC oxidation products. Those pathways are shown to
418 be important for SOA formation by modeling, field and chamber studies (Hodzic et al.,
419 2010; Jimenez et al., 2009; Murphy et al., 2017; Robinson et al., 2007; Shrivastava et al.,
420 2008; Tkacik et al., 2012; Zhao et al., 2014; Zhao et al., 2016a).

421 A sensitivity test was performed by using the newest CMAQ model version 5.3.1 that
422 includes all the above processes in the aerosol module. The SOA/POA ratio in Beijing is

423 improved greatly to be 0.83 in winter. However, high uncertainties still exist in the
424 emissions of the involved precursors and characterization of SOA formation through these
425 processes, needing further constrains by observations. Their influences on water
426 partitioning into OPM and ~~insignificant improvements~~non-ideality of the modeled
427 ~~OA~~organic-water mixture on SOA will be evaluated in S3-a future study.

428 Due to the lack of observed OC and OA in July of 2013, as an alternative, model
429 performances are evaluated by comparing predicted and observed PM_{2.5} at ground sites
430 (Figure S2 shows the comparison of modeled and observed PM_{2.5} at monitoring sites S1)
431 as shown in Figure S1 (a) ~~during July of 2013.~~S3. Generally, ~~our~~the model can well
432 reproduce the diurnal variation of PM_{2.5} in most regions. Predicted PM_{2.5} on ~~high~~
433 ~~concentration~~highly concentrated days ~~are~~is biased low ~~compared to observations,~~
434 especially in the North ~~Central~~China Plain (NCP). ~~The~~NCP region has the highest PM_{2.5}
435 ranging from 60 $\mu\text{g m}^{-3}$ to 300 $\mu\text{g m}^{-3}$ ~~compared to other regions.~~
436 PM_{2.5} is significant in cities in ~~the~~Northwest. This might be due to missing dust emissions
437 in the current inventory (Hu et al., 2016). To further evaluate the model performance,
438 ~~statistics of averaged~~MFB and MFE were of modeled PM_{2.5} are plotted against ~~observed~~
439 ~~PM_{2.5} concentration at all monitoring sites (observations of each site as shown in~~Figure
440 ~~S3).~~S4. The criteria of MFB and ~~goal~~MFE followed recommendations ~~of~~by Boylan and
441 Russell ~~(2006)~~(2006). Our model ~~performed~~performs well assince most of the predictions
442 meet the criteria and a large fraction (>58%) meet the goal. The averaged MFB and MFE
443 across all the sites are -0.28 and 0.39 ~~respectively~~, indicating slightly underestimate of
444 PM_{2.5} by the model.

445 **3.2 Impacts ~~of water partitioning~~ on SOA and ALW_{org}**

446 ~~Distribution~~The spatial distribution of SOA ~~varied~~varies greatly in the two seasons. In
447 winter, SOA is relatively high in eastern SCB and ~~in the~~ contiguous areas central and eastern
448 provinces of Shandong, Henan, Anhui, and Hubei ~~provinces~~ (Figure 2 and Figure ~~S4~~S5).
449 Monthly-averaged SOA concentrations in the above ~~two~~ areas are up to 25 and 15-20 μg

450 m^{-3} , respectively. ~~The major precursors of SOA Anthropogenic emissions are originated~~
451 ~~from anthropogenic the major sources of SOA (Figure S6),~~ such as dicarbonyl products of
452 ~~aromatics from the oxidation, xylenes of xylene~~ and toluene (Figure S5). (Hu et al., 2017).
453 In summer, surface SOA is high in ~~NE Northeast~~, NCP, and YRD ~~regions. Shanghai,~~
454 ~~Jiangsu province and coastal areas of Yellow Sea show the .~~ The highest SOA of $\sim 9\text{-}12 \mu\text{g}$
455 m^{-3} ~~occurs in Shanghai and Jiangsu provinces as well as the coastal area of the Yellow Sea,~~
456 ~~with the value of $\sim 9\text{-}16 \mu\text{g m}^{-3}$ at the surface and $\sim 20\text{-}25 \text{ mg m}^{-2}$ as in~~ the column total (col-
457 SOA) ~~in~~ of the atmosphere below 21 km (Figure S4S5). Different from winter SOA, a
458 significant fraction of summer SOA is originated from biogenic emissions ~~in Shanghai and~~
459 ~~Jiangsu province~~ (Figure S5S7). Anthropogenic SOA is ~~high highly concentrated in July~~
460 ~~in~~ the coastal areas of ~~the~~ Yellow Sea and Bohai Bay.

461 Combined effects of water partitioning into OPM and non-ideality on SOA formation
462 (S3) ~~also exhibit strong seasonal variation. In winter, the increase of SOA is relatively~~
463 ~~small, slightly decreased~~ by $\sim 1\text{-}4.5 \mu\text{g m}^{-3}$ (10-20%) ~~on average~~ at the surface (Figure 2)
464 and less than $\sim 51 \text{ mg m}^{-2}$ (~~10-30%~~) ~~as for 20%~~ in the column ~~concentration~~ (Figure S4).
465 ~~The influences on SOA also differ in different altitudes. For example, the maximum~~
466 ~~increment at the surface is observed in Shandong province in NCP (Figure 2), while SOA~~
467 ~~at higher levels of the atmosphere is more significant in South China (Figure S4). The~~
468 ~~increase in SOA is mostly attributed to anthropogenic sources in winter (Figure S5 and~~
469 ~~S7).)~~ ~~over high SOA regions where anthropogenic sources dominate. We show later that~~
470 ~~the decrease of SOA is mainly due to the large activity coefficients of SVOCs which~~
471 ~~decrease $K_{p,i}$.~~ In summer, higher temperature and ~~relative humidity (RH)~~ promote ~~water~~
472 ~~partitioning and~~ SOA formation ~~as well as water partitioning into OPM. At the surface, so~~
473 ~~that~~ SOA increases by $3\text{-}9 \mu\text{g m}^{-3}$ (40-50%) in coastal areas and $2\text{-}9 \mu\text{g m}^{-3}$ (20-90%) over
474 ~~the land, which are dominated by anthropogenic and biogenic origin, respectively (Figure~~
475 ~~S6). For col SOA, in addition to coastal areas, more significant increase is observed in~~
476 ~~YRD, most of Henan province, and the contiguous areas of Hubei, Hunan, and Jiangxi~~

477 province (Figure S4) by about 30-70% apparently over the entire domain, with the highest
478 enhancement of 2-4 $\mu\text{g m}^{-3}$ (20-50%) at the surface (Figure 2) and 4-6 mg m^{-2} (30-60%) in
479 the column (Figure S5) over YRD and the coastal area of Yellow Sea. Anthropogenic SOA
480 dominates the total change in winter as shown in Figure S6. In summer, the increase of
481 SOA is attributed to biogenic sources in inland areas and anthropogenic sources over the
482 ocean (Figure S7).

483 Regional distribution of ~~water partitioning into OPM~~ ALW_{org} is similar to the
484 ~~changes~~ of SOA, as shown in Figure 3 shows the regional distribution of monthly.
485 In winter, a maximum averaged daily maximum ALW_{org} . We see up to 9 concentration of
486 3.0 $\mu\text{g m}^{-3}$ for ALW_{org} at surface occurs in Shandong in winter where great increment in the
487 high SOA appears as well. In region, where significant changes of SOA also occur. In other
488 areas, the averaged concentration of ALW_{org} is about 2-6 0.5-1.5 $\mu\text{g m}^{-3}$. The Overall, the
489 average ratio of ALW_{org} to SOA is about 0.1-0.53 in winter. In summer, water partitioning
490 into OPM mostly involves occurs in the east coastal areas area at the surface where a
491 significant increase of anthropogenic SOA (such as those from toluene and xylenes) is
492 observed. This might be due to the high polarity of anthropogenic SVOCs (having more -
493 COOH groups) that absorb more water. In the coastal areas, ALW_{org} is about 5-12 $\mu\text{g m}^{-3}$,
494 with a ratio to SOA area, the averaged concentration of 0.3-0.6. ALW_{org} over the land is
495 about 25-7 $\mu\text{g m}^{-3}$, with the ALW_{org} /SOA ratio of 0.5-1.0. Over the land, the averaged
496 concentration of ALW_{org} is about 1-3 $\mu\text{g m}^{-3}$ (ALW_{org} /SOA ratio of 0.12-0.45) in most
497 areas, which Northeast and East China. Water partitioning is mostly associated with the
498 increase of BSOA such as biogenic SOA originated from isoprene and monoterpenes
499 oxidation that produces SVOCs with abundant OH group in SVOCs. The highest.

500 Based on the column concentrations of ALW_{org} is 16 $\mu\text{g m}^{-3}$ near Shanghai (and
501 ALW_{org} /SOA ratio of 0.57). Water partitioning also varies at different altitudes (Figure S9).
502 In (Figure S8), in winter, more column water partitions into OPM (col ALW_{org}) in
503 Chongqing, Hunan, Guanxi, Guangdong and Guizhou province, with the ALW_{org} must have

504 occurred in the south and southwest regions at higher levels where a significant increase of
505 col-SOA occurs (Figure S5). The averaged col-ALW_{org}/col-SOA ratio of in the high SOA
506 area is 0.21-0.3. In summer, the ALW_{org} must be high at higher altitudes over the central
507 regions in China. The maximum col-ALW_{org} is predicted about 7 mg m⁻² over the land,
508 especially in YRD, with the col-ALW_{org}/col-SOA ratio of 0.1 about 0.3 over eastern China.

509 Figure 4 shows **3.3 Impacts on aerosol properties**

510 Since ALW_{org} is determined in S3, the correlation values of κ_{org} with can be
511 estimated from the modeled ALW_{org}, OA and RH using Eq 4. 9 representative cities were
512 selected to investigate the relationship of κ_{org} vs. O:C ratio and its seasonal variation as
513 shown in Figure S9 and S10. The results of all the cities in winter and summer are merged
514 and analyzed as shown in Figure 4. Pairs of κ_{org} and O:C data are grouped into 10 O:C
515 bins and the averaged κ_{org} in each bin are then calculated. Overall, the estimated O:C
516 ratio is within the range of 0.2-0.6. In summer, the oxidation state of OA shows different
517 degrees of enhancement compared to winter at most of the cities except Guangzhou, due
518 to increased contribution of SOA to total OA. The averaged κ_{avg} of OA in each O:C bin
519 falls in the range of 0.001-0.18. The averaged κ_{org} in each O:C bin is less than 0.1 in
520 winter, with the highest κ_{avg} (~0.3) at Beijing value in Guangzhou. As more ALW_{org} is
521 formed in summer, the averaged κ_{org} also increases greatly with the highest value of
522 0.35 in Beijing. The linear correlation between κ_{org} and O:C shows significant spatial
523 and seasonal variations. For example, the slope of κ_{org} -O:C is much 70-90% smaller in
524 winter (45-74% less) than in summer in the Northern cities such as Shenyang, Beijing,
525 Zhengzhou, and Xi'an, while. However, in Guangzhou, the slope of κ_{avg} -O:C in winter is
526 much is 83% higher (47-104% more) in winter than in summer in the Southern cities, such
527 as Nanjing, . In Chengdu and Guangzhou. In Jinan and Shanghai, the slope is quite similar
528 in both seasons. The fitted correlations are very different from previous studies with a
529 relatively higher Overall, the slope of κ_{org} -O:C from vs. O:C in the 9 cities is 0.16 in winter
530 and 0.40 in summer. Most of the fitted linear correlations of the individual city fall outside

of the range of 0.18 to 0.37 suggested in previous studies (Duplissy et al., 2011; Lambe et al., 2011; Massoli et al., 2010; Chang et al., 2010) (Duplissy et al., 2011; Lambe et al., 2011; Massoli et al., 2010; Chang et al., 2010), indicating that the hygroscopicity of organic aerosols with chemical complexity cannot be simply represented by a single parameter such as the O:C ratio (Rickards et al., 2013) (Rickards et al., 2013).

3.3 Impacts on solar radiation

. In both seasons, κ_{org} approaches zero and negative values as O:C decreases, which might be due to the linear regression of κ_{org} and O:C. To avoid this, an exponential fitting of the two variables is performed so that κ_{org} falls in the range of (0,1) and is positively correlated with O:C. In this case, the fitted correlations are $\kappa_{org}=1-\exp(-(O:C/1.88)^{2.29})$ and $\kappa_{org}=1-\exp(-(O:C/1.06)^{4.50})$ for January and July of 2013, respectively.

The impacts on aerosol optical depth (AOD) and aerosol radiative forcing (ARF) were further investigated. Figure 5 shows the monthly-averaged AOD at 550 nm in January and July of 2013. It was calculated as by summarizing the accumulation product of model estimated extinction coefficient of fine particles ($EXT_i b_{ext,i}$) multiplied by the thickness (HL_i) in each layer:

$$AOD = \sum_{i=1}^N EXT_i \sum_{i=1}^N b_{ext,i} \times HL_i \quad (Eq6)$$

Where N is the number of layers. There are two methods to estimate the aerosol extinction coefficient in CMAQv5.0.1. One is using based on the Mie theory (EXT_m and the predicted aerosol component concentrations ($b_{ext,m}$), and the other is based on extinction values correlation from the IMPROVE monitoring network that considers the impacts of hygroscopicity of different aerosol components (EXT_r) ($b_{ext,r}$) (Malm et al., 1994) (Malm et al., 1994). AOD calculated with the two types of extinction coefficient are denoted as AOD_m and AOD_r, respectively.

In Figure 5, a clear pattern of high AOD_r in SCB and NCP and low AOD_r in west China is observed in both winter and summer, consistent with previous studies

556 ~~(He et al., 2019; He et al., 2016; Luo et al., 2014)~~(He et al., 2019; He et al., 2016; Luo et al.,
557 2014). An ~~identified~~identical trend in AOD_m is ~~observed as~~ shown in Figure ~~S10~~S11. The
558 monthly-averaged AOD_r ranges from ~~1.40~~ to ~~3.52~~ in January and from ~~0.43~~ to ~~0.89~~ in July.
559 AOD_m is lower than AOD_r, falling in 0.7-2.2 in January and 0.3-0.6 in July. The model
560 significantly overestimates AOD in January but agrees better with observations from
561 MODIS ~~in the where AOD is high regions~~ in July (Figure ~~S11~~S12). ~~The bias~~ in the
562 predicted AOD might be partially due to the empirical equation applied in the calculation
563 of AOD in CMAQ (~~Wang et al., 2009; Liu et al., 2010~~)(Wang et al., 2009; Liu et al., 2010),
564 and partially due to the uncertainties of fine AOD overland from MODIS data (~~Wang et al.,~~
565 ~~2009; Levy et al., 2010~~)(Wang et al., 2009; Levy et al., 2010). ~~With water partitioning into~~
566 ~~OPM, changes in SOA mass concentration and chemical composition lead to~~The increase
567 of AOD, ~~which due to ALW_{org}~~ shows a strong spatial and seasonal pattern. In winter, there
568 ~~is~~are no significant ~~increase~~changes in AOD_r across the whole domain, due to insignificant
569 changes of SOA. In summer, AOD_r increases significantly in YRD and the adjacent ~~area~~
570 ~~of Hubei, Hunan, and Jiangxi province~~areas by up to ~~42~~10%.

571 ARF represents the ~~changes~~change in the radiative flux at the top of the atmosphere
572 due to aerosols. ~~The off-line~~An offline version of the Shortwave Radiative Transfer Model
573 For GCMs (RRTMG_SW) ~~is~~was used to calculate the direct radiative effect of aerosols on
574 shortwave radiation (~~Iacono et al., 2008~~)(Iacono et al., 2008). Generally, fine aerosols
575 exhibit cooling effects on the shortwave radiation in both winter and summer over the entire
576 domain as shown in Figure 6. This impact is much stronger in the areas where AOD is high
577 (Figure 5). The ARF ~~at top of atmosphere (TOA)~~ is highest in Shandong in winter and in
578 the coastal areas near Jiangsu province in summer, which are about ~~-125~~ W m⁻² and -
579 ~~96~~ W m⁻², respectively. In winter, no significant changes of ARF are observed in ~~the high~~
580 ~~regions of~~ eastern China (Figure ~~6b~~6b). This is likely attributed to an insignificant
581 contribution of SOA to PM_{2.5} in winter compared to other components with cooling effects,
582 such as sulfate. In summer, SOA is an important component of PM_{2.5} (20-60%), and the

583 effects of water partitioning on shortwave radiation ~~is~~are relatively stronger. An
584 enhancement of up to 15% in the cooling effects of ARF occurs near ~~the~~ YRD region where
585 AOD significantly changes as well.

586 **4 Discussion**

587 **3.4 Sensitivity to T, RH, and SVP**

588 Meteorological conditions and SOA precursors affect the impacts of water partitioning on
589 SOA. ~~Figure 7 shows the effects of different factors on the daily maximum change of SOA~~
590 ~~in Jinan and Nanjing, two representative cities in winter and summer, respectively. As~~
591 ~~shown in Figure 7(a), the daily maximum elevation of SOA occurs when RH is greater than~~
592 ~~70% in both cities. This is consistent with the previous study in the Southeast U.S. during~~
593 ~~summer (Pankow et al., 2015). A clear correlation of the changes in SOA with SOA~~
594 ~~concentration in Nanjing (R=0.84) during summer can be observed. However, this~~
595 ~~correlation is relatively weak in Jinan (R=0.44) during winter. There is no strong~~
596 ~~correlation between changes in SOA and temperature as shown in Figure 7(b), likely due~~
597 ~~to the daily variation of SOA mass and composition.~~ To better illustrate the dependency of
598 SOA on temperature, RH, and ~~relative humidity~~SVP of SVOCs, an offline calculation of
599 SOA formation was performed at two representative cities (Jinan in NCP during winter
600 and Nanjing in YRD during summer) when the daily maximum SOA increases occurred.
601 We ~~assumed~~assume temperature (T) and water vapor mixing ratio (QV) to be within the
602 range of $\bar{X} \pm \sigma$, where \bar{X} and σ are the mean and standard deviation calculated based on
603 WRF ~~prediction~~predictions at each location. We ~~chose~~choose 10 evenly distributed values
604 for T and QV within the range of $\bar{X} \pm \sigma$. The temperature dependence parameter of
605 ~~saturation vapor pressure~~SVP (ΔH) ~~was~~/R is also scaled by 0.2, 0.8, 1.4, and 2.0 separately
606 for all the SVOCs. As shown in Figure ~~8~~7, SOA ~~indicates~~exhibits a negative correlation
607 with ~~temperature~~T and a positive correlation with ~~RH~~QV in both cities. SOA is more
608 sensitive to ~~RH~~QV under ~~cool~~cold conditions (~~JN~~Jinan) and to temperature under hot
609 conditions (~~NJ~~). ~~An interesting finding is that significant increases in SOA Nanjing). When~~

610 ~~the temperature is fixed, the sensitivity of SOA to $\Delta H/R$ is different~~ in the two cities ~~occur~~
611 ~~during different time periods of the day. Water partitioning tends to affect SOA in the~~
612 ~~afternoon and evening in Jinan, which mostly happens in the early morning and at noon in~~
613 ~~Nanjing. The different timing is likely. We find more changes in SOA across $\Delta H/R$ in~~
614 ~~Jinan. This is attributed to a substantial increase in SOA precursors in the two cities. In~~
615 ~~Jinan, the most contributing SVOCs are originated~~ the temperature correction factor (ζ_{corr})
616 of K_p in CMAQ as defined below:

$$\zeta_{corr} = \frac{K_{p,T_{ref}}}{K_{p,T}} = \frac{T_{ref}}{T} \exp \left[\frac{\Delta H}{R} \left(\frac{1}{T_{ref}} - \frac{1}{T} \right) \right] \quad (\text{Eq8})$$

617 where T_{ref} is the reference temperature (298K) and T is temperature. According to
618 ~~Figure 7, the range of T is 265-274K in Jinan and 300-307K in Nanjing. The deviation of~~
619 ~~temperature from toluene and xylenes oxidation, as well as oligomers formed by their~~
620 ~~oxidation products in OPM. Possible emission sources include transportation, petroleum~~
621 ~~refining, manufacturing, painting, etc. SOA increase in Nanjing is mostly associated with~~
622 ~~biogenic sources including isoprene and monoterpenes.~~ the reference value (298K) is
623 greater in Jinan than in Nanjing. Therefore, the unit change of $\Delta H/R$ causes greater
624 variations of ζ_{corr} and thus K_p in Jinan. As a result, SOA is more sensitive to $\Delta H/R$.
625 The impacts of SVP estimation on SOA are more significant in winter than in summer.

626 **3.5 Separate impacts of water partitioning/polarity of condensed organics**

627 Impacts of water partitioning into OPM and non-ideality of ~~the~~ organic-water mixture on
628 SOA are in opposite directions. Water partitioning ~~alone~~ increases SOA by ~~~10-20-60%~~ in
629 winter and ~~~20-100~~ 30-80% in summer in most areas of the domain (Figure ~~S128~~). This is
630 because ~~that~~ the molecular weight of water is quite small and will reduce the molar
631 averaged weight of OPM (\overline{MW}) in Eq 2 (Pankow et al., 2015) (Pankow et al., 2015).
632 The reduced \overline{MW} further increases ~~$K_{p,i}$~~ K_p and promotes mass transfer of
633 SVOCs from the gas-phase to ~~the~~ OPM. On the other hand, by considering non-ideality of
634 ~~the~~ organic-water mixture, activity coefficients of SVOCs are usually greater than 1.0 in

635 this study, leading to a decrease in $K_{p,i}K_p$. As a result, the total SOA concentration is
636 reduced by ~~up to~~ ~10-20% in winter and ~~~30~~10-50% in summer ~~in the high regions~~ (Figure
637 ~~S139~~). Overall, the final impacts are the combined consequences of the two “processes”.
638 In winter, the increase of SOA caused by water partitioning is offset by the decrease of
639 SOA due to the polarity of SVOCs in most areas of the domain, resulting in ~~no significant~~
640 ~~changes.slight decreases of SOA~~. In summer, ~~effects~~the effect of water partitioning
641 ~~overcome~~overcomes that of SVOC polarity so ~~as~~ the total SOA loading increases. This
642 further leads to an enhanced attenuation of shortwave solar radiation and cooling of the
643 atmosphere.

644 ~~5 Conclusion~~

645 4 Conclusions

646 The WRF/CMAQ model was used to investigate the impacts of water partitioning into
647 OPM and non-ideality of the organic-water mixture on SOA formation over eastern China
648 during January and July of 2013. SOA is greatly enhanced in summer especially in YRD
649 and over the Yellow Sea by up to ~~90~~50% and ~~70~~60% at the surface and in the whole column,
650 respectively. No significant impacts on SOA are observed in winter. This might be due to
651 the underestimation of SOA in the current model. ALW_{org} is highly correlated with the
652 ~~changes~~change of SOA, with the ratio of ALW_{org} to SOA of 0.1-0.53 and 0.2-1-0.6 at the
653 surface where significant changes of SOA occur in winter and summer, respectively. By
654 using the modeled ALW_{org} , correlations between κ_{org} and O:C ~~were~~are examined in 9
655 representative cities, showing significant spatial and seasonal variations. The increases in
656 SOA lead to ~~12% elevation of AOD and 15% an~~ enhancement in the averaged AOD and
657 the cooling effects of ARFaerosols, by up to 10% and 15% respectively in summer. The
658 model predicted SOA is sensitive to temperature and QV in both seasons, with higher
659 sensitivity to QV during winter and temperature during summer. Estimation of SVP also
660 affects modeled SOA, especially in a cold environment. The effects of water partitioning
661 into OPM and non-ideality of the organic-water mixture on SOA ~~were also examined~~

662 ~~separately~~ are the opposite. Since the activity coefficients of SVOCs are mostly greater than
663 1.0 during the simulated episode, SOA ~~concentrations decrease~~ concentration decreases
664 when the non-ideality effect is considered. Daily Averaged SOA concentration decreases
665 by up to ~~~1020%~~ in winter and ~~~3050%~~ in summer ~~in the high regions~~. Water partitioning
666 alone increases SOA by ~~~10-20-60%~~ in winter and ~~~20-10030-80%~~ in summer. It should
667 be noticed that the results shown in this study are the lower limit as the current model tends
668 to underestimate SOA. It is crucial to consider both effects in simulating SOA formation
669 under hot and humid conditions in CTMs.

671 **Data availability**

672 Data used in this manuscript can be provided upon request by e-mail to the corresponding
673 authors, Qi Ying (qying@civil.tamu.edu) and Jianlin Hu (jianlinhu@nuist.edu.cn).

675 **Competing interests**

676 The authors declare that they have no conflict of interest.

678 **Acknowledgment**

679 This project was partly supported by the National Key R&D Program of China
680 (2018YFC0213802, Task #2), National Science Foundation of China (No. 41705102 and
681 41875149), and the Major Research Plan of National Social Science Foundation
682 (18ZDA052). The authors thank James F. Pankow for providing the SOA module code.
683 Jingyi Li acknowledges support from the Startup Foundation for Introducing Talent of
684 NUIST grant no. 2243141701014, the Priority Academic Program Development of Jiangsu
685 Higher Education Institutions (PAPD), and Six Talent Peaks Project of Jiangsu Province.

687 **References**

688 Ansari, A. S., and Pandis, S. N.: Water Absorption by Secondary Organic Aerosol and Its
689 Effect on Inorganic Aerosol Behavior, Environ. Sci. Technol., 34, 71-77,
690 10.1021/es990717q, 2000.
691 Asa-Awuku, A., Nenes, A., Gao, S., Flagan, R. C., and Seinfeld, J. H.: Water-soluble SOA
692 from Alkene ozonolysis: composition and droplet activation kinetics inferences from

693 analysis of CCN activity, *Atmos. Chem. Phys.*, 10, 1585-1597, 10.5194/acp-10-1585-2010,
694 2010.

695 Atkinson, R. W., Kang, S., Anderson, H. R., Mills, I. C., and Walton, H. A.:
696 Epidemiological time series studies of PM_{2.5} and daily mortality and hospital admissions:
697 a systematic review and meta-analysis, *Thorax*, 69, 660-665, 10.1136/thoraxjnl-2013-
698 204492 %J Thorax, 2014.

699 Ayers, G. P.: Comment on regression analysis of air quality data, *Atmos. Environ.*, 35,
700 2423-2425, [https://doi.org/10.1016/S1352-2310\(00\)00527-6](https://doi.org/10.1016/S1352-2310(00)00527-6), 2001.

701 [Baek, J., Hu, Y., Odman, M. T., and Russell, A. G.: Modeling secondary organic aerosol in CMAQ using](#)
702 [multigenerational oxidation of semi-volatile organic compounds, 116, 10.1029/2011jd015911, 2011.](#)

703 Bergström, R., Denier van der Gon, H. A. C., Prévôt, A. S. H., Yttri, K. E., and Simpson,
704 D.: Modelling of organic aerosols over Europe (2002–2007) using a volatility basis
705 set (VBS) framework: application of different assumptions regarding the formation of
706 secondary organic aerosol, *Atmos. Chem. Phys.*, 12, 8499-8527, 10.5194/acp-12-8499-
707 2012, 2012.

708 Bowman, F. M., and Melton, J. A.: Effect of activity coefficient models on predictions of
709 secondary organic aerosol partitioning, *J. Aerosol Sci.*, 35, 1415-1438,
710 <https://doi.org/10.1016/j.jaerosci.2004.07.001>, 2004.

711 Boylan, J. W., and Russell, A. G.: PM and light extinction model performance metrics,
712 goals, and criteria for three-dimensional air quality models, *Atmos. Environ.*, 40, 4946-
713 4959, <https://doi.org/10.1016/j.atmosenv.2005.09.087>, 2006.

714 [Budisulistiorini, S. H., Nenes, A., Carlton, A. G., Surratt, J. D., McNeill, V. F., and Pye,](#)
715 [H. O. T.: Simulating Aqueous-Phase Isoprene-Epoxydiol \(IEPOX\) Secondary Organic](#)
716 [Aerosol Production During the 2013 Southern Oxidant and Aerosol Study \(SOAS\),](#)
717 [Environ. Sci. Technol., 51, 5026-5034, 10.1021/acs.est.6b05750, 2017.](#)

718 Cappa, C. D., Lovejoy, E. R., and Ravishankara, A. R.: Evidence for liquid-like and
719 nonideal behavior of a mixture of organic aerosol components, *P. Natl. Acad. Sci. USA*,
720 105, 18687-18691, 10.1073/pnas.0802144105, 2008.

721 [Carlton, A. G., Bhawe, P. V., Napelenok, S. L., Edney, E. O., Sarwar, G., Pinder, R. W.,](#)
722 [Pouliot, G. A., and Houyoux, M.: Model Representation of Secondary Organic Aerosol in](#)
723 [CMAQv4.7, Environ. Sci. Technol., 44, 8553-8560, 10.1021/es100636q, 2010.](#)

724 Chang, R. Y. W., Slowik, J. G., Shantz, N. C., Vlasenko, A., Liggió, J., Sjostedt, S. J.,
725 Leaitch, W. R., and Abbatt, J. P. D.: The hygroscopicity parameter (κ) of ambient organic
726 aerosol at a field site subject to biogenic and anthropogenic influences: relationship to
727 degree of aerosol oxidation, *Atmos. Chem. Phys.*, 10, 5047-5064, 10.5194/acp-10-5047-
728 2010, 2010.

729 Denjean, C., Formenti, P., Picquet-Varrault, B., Pangui, E., Zapf, P., Katrib, Y., Giorio, C.,
730 Tapparo, A., Monod, A., Temime-Roussel, B., Decorse, P., Mangeney, C., and Doussin, J.
731 F.: Relating hygroscopicity and optical properties to chemical composition and structure
732 of secondary organic aerosol particles generated from the ozonolysis of α -pinene, *Atmos.*
733 *Chem. Phys.*, 15, 3339-3358, 10.5194/acp-15-3339-2015, 2015.

734 Donahue, N. M., Robinson, A. L., Stanier, C. O., and Pandis, S. N.: Coupled partitioning,
735 dilution, and chemical aging of semivolatile organics, *Environ. Sci. Technol.*, 40, 02635-
736 02643, 10.1021/es052297c, 2006.

737 Duplissy, J., DeCarlo, P. F., Dommen, J., Alfarra, M. R., Metzger, A., Barmapadimos, I.,
738 Prevot, A. S. H., Weingartner, E., Tritscher, T., Gysel, M., Aiken, A. C., Jimenez, J. L.,
739 Canagaratna, M. R., Worsnop, D. R., Collins, D. R., Tomlinson, J., and Baltensperger, U.:
740 Relating hygroscopicity and composition of organic aerosol particulate matter, *Atmos.*
741 *Chem. Phys.*, 11, 1155-1165, 10.5194/acp-11-1155-2011, 2011.

742 Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M.,
743 Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M.,
744 Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T.,
745 Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T.,
746 Petäjä, T., Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop, D. R., Wildt, J., and
747 Mentel, T. F.: A large source of low-volatility secondary organic aerosol, *Nature*, 506, 476,
748 10.1038/nature13032, 2014.

749 [El-Sayed, M. M. H., Ortiz-Montalvo, D. L., and Hennigan, C. J.: The effects of isoprene](#)
750 [and NOx on secondary organic aerosols formed through reversible and irreversible uptake](#)
751 [to aerosol water, *Atmos. Chem. Phys.*, 18, 1171-1184, 10.5194/acp-18-1171-2018, 2018.](#)

752 [Emery, C., Tai, E., and Yarwood, G.: Enhanced meteorological modeling and performance](#)
753 [evaluation for two texas episodes, Report to the Texas Natural Resources Conservation](#)
754 [Commission, prepared by ENVIRON, International Corp., Novato, CA., available at:](#)
755 [http://www.tceq.state.tx.us/assets/public/implementation/air/am/contracts/reports/mm/En](http://www.tceq.state.tx.us/assets/public/implementation/air/am/contracts/reports/mm/EnhancedMetModelingAndPerformanceEvaluation.pdf)
756 [hancedMetModelingAndPerformanceEvaluation.pdf](http://www.tceq.state.tx.us/assets/public/implementation/air/am/contracts/reports/mm/EnhancedMetModelingAndPerformanceEvaluation.pdf), 2001.

757 [EPA: U.S.: Guidance on the Use of Models and Other Analyses for Demonstrating](#)
758 [Attainment of Air Quality Goals for Ozone, PM2.5, and Regional Haze, EPA-454/B-07-](#)
759 [002, 2007.](#)

760 Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in cloud
761 droplets and aqueous particles (aqSOA): a review of laboratory, field and model studies,
762 *Atmos. Chem. Phys.*, 11, 11069-11102, 10.5194/acp-11-11069-2011, 2011.

763 [Feng, S., Gao, D., Liao, F., Zhou, F., and Wang, X.: The health effects of ambient PM2.5 and potential](#)
764 [mechanisms, *Ecotox. Environ. Safe.*, 128, 67-74, <https://doi.org/10.1016/j.ecoenv.2016.01.030>, 2016.](#)

765 Fredenslund, A., Jones, R. L., and Prausnitz, J. M.: Group-contribution estimation of
766 activity coefficients in nonideal liquid mixtures, *AIChE J.*, 21, 1086-1099,
767 doi:10.1002/aic.690210607, 1975.

768 Fu, H., and Chen, J.: Formation, features and controlling strategies of severe haze-fog
769 pollutions in China, *Sci. Total Environ.*, 578, 121-138,
770 <https://doi.org/10.1016/j.scitotenv.2016.10.201>, 2017.

771 [Galloway, M. M., Chhabra, P. S., Chan, A. W. H., Surratt, J. D., Flagan, R. C., Seinfeld, J.](#)
772 [H., and Keutsch, F. N.: Glyoxal uptake on ammonium sulphate seed aerosol: reaction](#)
773 [products and reversibility of uptake under dark and irradiated conditions, *Atmos. Chem.*](#)
774 [Phys., 9, 3331-3345, 10.5194/acp-9-3331-2009, 2009.](#)

775 Gentner, D. R., Jathar, S. H., Gordon, T. D., Bahreini, R., Day, D. A., El Haddad, I., Hayes,
776 P. L., Pieber, S. M., Platt, S. M., de Gouw, J., Goldstein, A. H., Harley, R. A., Jimenez, J.
777 L., Prévôt, A. S. H., and Robinson, A. L.: Review of Urban Secondary Organic Aerosol
778 Formation from Gasoline and Diesel Motor Vehicle Emissions, *Environ. Sci. Technol.*, 51,
779 1074-1093, 10.1021/acs.est.6b04509, 2017.

780 Guo, H., Xu, L., Bougiatioti, A., Cerully, K. M., Capps, S. L., Hite Jr, J. R., Carlton, A. G.,
781 Lee, S. H., Bergin, M. H., Ng, N. L., Nenes, A., and Weber, R. J.: Fine-particle water and
782 pH in the southeastern United States, *Atmos. Chem. Phys.*, 15, 5211-5228, 10.5194/acp-
783 15-5211-2015, 2015.

~~784 Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M., Zeng, L., Molina,
785 M. J., and Zhang, R.: Elucidating severe urban haze formation in China, *P. Natl. Acad. Sci. USA*, 111,
786 17373-17378, 10.1073/pnas.1419604111, 2014.~~

~~787 Hatzianastassiou, N., Matsoukas, C., Drakakis, E., Stackhouse Jr, P. W., Koepke, P., Fotiadis, A., Pavlakis,
788 K. G., and Vardavas, I.: The direct effect of aerosols on solar radiation based on satellite observations,
789 reanalysis datasets, and spectral aerosol optical properties from Global Aerosol Data Set (GADS),
790 *Atmos. Chem. Phys.*, 7, 2585-2599, 10.5194/acp-7-2585-2007, 2007.~~

791 Hayes, P. L., Carlton, A. G., Baker, K. R., Ahmadov, R., Washenfelder, R. A., Alvarez, S.,
792 Rappenglück, B., Gilman, J. B., Kuster, W. C., de Gouw, J. A., Zotter, P., Prévôt, A. S. H.,
793 Szidat, S., Kleindienst, T. E., Offenberg, J. H., Ma, P. K., and Jimenez, J. L.: Modeling the
794 formation and aging of secondary organic aerosols in Los Angeles during CalNex 2010,
795 *Atmos. Chem. Phys.*, 15, 5773-5801, 10.5194/acp-15-5773-2015, 2015.

796 He, Q., Zhang, M., and Huang, B.: Spatio-temporal variation and impact factors analysis
797 of satellite-based aerosol optical depth over China from 2002 to 2015, *Atmos. Environ.*,
798 129, 79-90, <https://doi.org/10.1016/j.atmosenv.2016.01.002>, 2016.

799 He, Q., Gu, Y., and Zhang, M.: Spatiotemporal patterns of aerosol optical depth throughout
800 China from 2003 to 2016, *Sci. Total Environ.*, 653, 23-35,
801 <https://doi.org/10.1016/j.scitotenv.2018.10.307>, 2019.

~~802 Heald, C. L., Jacob, D. J., Turquety, S., Hudman, R. C., Weber, R. J., Sullivan, A. P., Peltier, R. E., Atlas, E.
803 L., de Gouw, J. A., Warneke, C., Holloway, J. S., Neuman, J. A., Flocke, F. M., and Seinfeld, J. H.:
804 Concentrations and sources of organic carbon aerosols in the free troposphere over North America,
805 *J. Geophys. Res.*, 111, D23S47, 10.1029/2006jd007705, 2006.~~

806 Healy, R. M., Temime, B., Kuprovskite, K., and Wenger, J. C.: Effect of Relative
807 Humidity on Gas/Particle Partitioning and Aerosol Mass Yield in the Photooxidation of p-
808 Xylene, *Environ. Sci. Technol.*, 43, 1884-1889, 10.1021/es802404z, 2009.

~~809 Hegerl, G. C., Black, E., Allan, R. P., Ingram, W. J., Polson, D., Trenberth, K. E., Chadwick, R. S., Arkin, P.
810 A., Sarojini, B. B., Becker, A., Dai, A., Durack, P. J., Easterling, D., Fowler, H. J., Kendon, E. J., Huffman, G.
811 J., Liu, C., Marsh, R., New, M., Osborn, T. J., Skliris, N., Stott, P. A., Vidale, P. L., Wijffels, S. E., Wilcox, L.
812 J., Willett, K. M., and Zhang, X.: CHALLENGES IN QUANTIFYING CHANGES IN THE GLOBAL WATER
813 CYCLE, *Bull. Am. Meteorol. Soc.*, 96, 1097-1116, 2015.~~

~~814 Hong, J., Xu, H., Tan, H., Yin, C., Hao, L., Li, F., Cai, M., Deng, X., Wang, N., Su, H., Cheng, Y., Wang, L.,
815 Petäjä, T., and Kerminen, V. M.: Mixing state and particle hygroscopicity of organic-dominated aerosols~~

816 over the Pearl River Delta region in China, *Atmos. Chem. Phys.*, **18**, 14079–14094, [10.5194/acp-18-](https://doi.org/10.5194/acp-18-14079-2018)
817 [14079-2018](https://doi.org/10.5194/acp-18-14079-2018), 2018.

818 Hodzic, A., Jimenez, J. L., Madronich, S., Canagaratna, M. R., DeCarlo, P. F., Kleinman,
819 L., and Fast, J.: Modeling organic aerosols in a megacity: potential contribution of semi-
820 volatile and intermediate volatility primary organic compounds to secondary organic
821 aerosol formation, *Atmos. Chem. Phys.*, **10**, 5491-5514, [10.5194/acp-10-5491-2010](https://doi.org/10.5194/acp-10-5491-2010), 2010.

822 Hu, J., Chen, J., Ying, Q., and Zhang, H.: One-year simulation of ozone and particulate
823 matter in China using WRF/CMAQ modeling system, *Atmos. Chem. Phys.*, **16**, 10333-
824 10350, [10.5194/acp-16-10333-2016](https://doi.org/10.5194/acp-16-10333-2016), 2016.

825 Hu, J., Wang, P., Ying, Q., Zhang, H., Chen, J., Ge, X., Li, X., Jiang, J., Wang, S., Zhang,
826 J., Zhao, Y., and Zhang, Y.: Modeling biogenic and anthropogenic secondary organic
827 aerosol in China, *Atmos. Chem. Phys.*, **17**, 77-92, [10.5194/acp-17-77-2017](https://doi.org/10.5194/acp-17-77-2017), 2017.

828 Huang, R.-J., Zhang, Y., Bozzetti, C., Ho, K.-F., Cao, J.-J., Han, Y., Daellenbach, K. R.,
829 Slowik, J. G., Platt, S. M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A.,
830 Crippa, M., Ciarelli, G., Piazzalunga, A., Schwikowski, M., Abbaszade, G., Schnelle-Kreis,
831 J., Zimmermann, R., An, Z., Szidat, S., Baltensperger, U., Haddad, I. E., and Prévôt, A. S.
832 H.: High secondary aerosol contribution to particulate pollution during haze events in
833 China, *Nature*, **514**, 218, [10.1038/nature13774](https://doi.org/10.1038/nature13774), 2014.

834 ~~Huang, W., Saathoff, H., Shen, X., Ramisetty, R., Leisner, T., and Mohr, C.: Chemical Characterization of~~
835 ~~Highly Functionalized Organonitrates Contributing to Night Time Organic Aerosol Mass Loadings and~~
836 ~~Particle Growth, *Environ. Sci. Technol.*, **53**, 1165–1174, [10.1021/acs.est.8b05826](https://doi.org/10.1021/acs.est.8b05826), 2019.~~

837 Iacono, M. J., Delamere, J. S., Mlawer, E. J., Shephard, M. W., Clough, S. A., and Collins,
838 W. D.: Radiative forcing by long-lived greenhouse gases: Calculations with the AER
839 radiative transfer models, *J. Geophys. Res.*, **113**, D13103, [10.1029/2008jd009944](https://doi.org/10.1029/2008jd009944), 2008.

840 Jathar, S. H., Mahmud, A., Barsanti, K. C., Asher, W. E., Pankow, J. F., and Kleeman, M.
841 J.: Water uptake by organic aerosol and its influence on gas/particle partitioning of
842 secondary organic aerosol in the United States, *Atmos. Environ.*, **129**, 142-154,
843 <https://doi.org/10.1016/j.atmosenv.2016.01.001>, 2016.

844 Jiang, F., Liu, Q., Huang, X., Wang, T., Zhuang, B., and Xie, M.: Regional modeling of
845 secondary organic aerosol over China using WRF/Chem, *J. Aerosol Sci.*, **43**, 57-73,
846 <https://doi.org/10.1016/j.jaerosci.2011.09.003>, 2012.

847 ~~Jokinen, T., Berndt, T., Makkonen, R., Kerminen, V. M., Junninen, H., Paasonen, P., Stratmann, F.,~~
848 ~~Herrmann, H., Guenther, A. B., Worsnop, D. R., Kulmala, M., Ehn, M., and Sipilä, M.: Production of~~
849 ~~extremely low volatile organic compounds from biogenic emissions: Measured yields and atmospheric~~
850 ~~implications, **112**, 7123–7128, [10.1073/pnas.1423977112](https://doi.org/10.1073/pnas.1423977112) %J Proceedings of the National Academy of~~
851 ~~Sciences, 2015.~~

852 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J.
853 H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich,
854 I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V.
855 A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J.,
856 Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J.,

857 [Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo,](#)
858 [Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D.,](#)
859 [Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y.,](#)
860 [Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C.,](#)
861 [Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E.,](#)
862 [Baltensperger, U., and Worsnop, D. R.: Evolution of Organic Aerosols in the Atmosphere,](#)
863 [Science, 326, 1525-1529, 10.1126/science.1180353, 2009.](#)

864 Kim, Y., Sartelet, K., and Couvidat, F.: Modeling the effect of non-ideality, dynamic mass
865 transfer and viscosity on SOA formation in a 3-D air quality model, *Atmos. Chem. Phys.*,
866 19, 1241-1261, 10.5194/acp-19-1241-2019, 2019.

867 [Knote, C., Hodzic, A., Jimenez, J. L., Volkamer, R., Orlando, J. J., Baidar, S., Brioude, J.,](#)
868 [Fast, J., Gentner, D. R., Goldstein, A. H., Hayes, P. L., Knighton, W. B., Oetjen, H., Setyan,](#)
869 [A., Stark, H., Thalman, R., Tyndall, G., Washenfelder, R., Waxman, E., and Zhang, Q.:](#)
870 [Simulation of semi-explicit mechanisms of SOA formation from glyoxal in aerosol in a 3-](#)
871 [D model, Atmos. Chem. Phys., 14, 6213-6239, 10.5194/acp-14-6213-2014, 2014.](#)

872 Kurokawa, J., Ohara, T., Morikawa, T., Hanayama, S., Janssens-Maenhout, G., Fukui, T.,
873 Kawashima, K., and Akimoto, H.: Emissions of air pollutants and greenhouse gases over
874 Asian regions during 2000–2008: Regional Emission inventory in ASia (REAS) version 2,
875 *Atmos. Chem. Phys.*, 13, 11019-11058, 10.5194/acp-13-11019-2013, 2013.

876 [Lai, S., Zhao, Y., Ding, A., Zhang, Y., Song, T., Zheng, J., Ho, K. F., Lee, S.-c., and Zhong,](#)
877 [L.: Characterization of PM_{2.5} and the major chemical components during a 1-year](#)
878 [campaign in rural Guangzhou, Southern China, Atmospheric Research, 167, 208-215,](#)
879 <https://doi.org/10.1016/j.atmosres.2015.08.007>, 2016.

880 Lambe, A. T., Onasch, T. B., Massoli, P., Croasdale, D. R., Wright, J. P., Ahern, A. T.,
881 Williams, L. R., Worsnop, D. R., Brune, W. H., and Davidovits, P.: Laboratory studies of
882 the chemical composition and cloud condensation nuclei (CCN) activity of secondary
883 organic aerosol (SOA) and oxidized primary organic aerosol (OPOA), *Atmos. Chem. Phys.*,
884 11, 8913-8928, 10.5194/acp-11-8913-2011, 2011.

885 Levy, R. C., Remer, L. A., Kleidman, R. G., Mattoo, S., Ichoku, C., Kahn, R., and Eck, T.
886 F.: Global evaluation of the Collection 5 MODIS dark-target aerosol products over land,
887 *Atmos. Chem. Phys.*, 10, 10399-10420, 10.5194/acp-10-10399-2010, 2010.

888 Li, J., Cleveland, M., Ziemba, L. D., Griffin, R. J., Barsanti, K. C., Pankow, J. F., and Ying,
889 Q.: Modeling regional secondary organic aerosol using the Master Chemical Mechanism,
890 *Atmos. Environ.*, 102, 52-61, <https://doi.org/10.1016/j.atmosenv.2014.11.054>, 2015.

891 Li, J., Zhang, M., Wu, F., Sun, Y., and Tang, G.: Assessment of the impacts of aromatic
892 VOC emissions and yields of SOA on SOA concentrations with the air quality model
893 RAMS-CMAQ, *Atmos. Environ.*, 158, 105-115,
894 <https://doi.org/10.1016/j.atmosenv.2017.03.035>, 2017a.

895 ~~Li, M., Zhang, Q., Streets, D. G., He, K. B., Cheng, Y. F., Emmons, L. K., Huo, H., Kang, S. C., Lu, Z., Shao,~~
896 ~~M., Su, H., Yu, X., and Zhang, Y.: Mapping Asian anthropogenic emissions of non-methane volatile~~
897 ~~organic compounds to multiple chemical mechanisms, Atmos. Chem. Phys., 14, 5617-5638,~~

898 ~~10.5194/acp-14-5617-2014, 2014.~~
899 ~~Li, X., Song, S., Zhou, W., Hao, J., Worsnop, D. R., and Jiang, J.: Interactions between~~
900 ~~aerosol organic components and liquid water content during haze episodes in Beijing,~~
901 ~~Atmos. Chem. Phys., 19, 12163-12174, 10.5194/acp-19-12163-2019, 2019.~~
902 ~~Li, Y. J., Sun, Y., Zhang, Q., Li, X., Li, M., Zhou, Z., and Chan, C. K.: Real-time chemical~~
903 ~~characterization of atmospheric particulate matter in China: A review, Atmos. Environ.,~~
904 ~~158, 270-304, <https://doi.org/10.1016/j.atmosenv.2017.02.027>, 2017b.~~
905 ~~[Lim, Y. B., Tan, Y., and Turpin, B. J.: Chemical insights, explicit chemistry, and yields of](#)~~
906 ~~[secondary organic aerosol from OH radical oxidation of methylglyoxal and glyoxal in the](#)~~
907 ~~[aqueous phase, Atmos. Chem. Phys., 13, 8651-8667, 10.5194/acp-13-8651-2013, 2013.](#)~~
908 ~~Lin, J., An, J., Qu, Y., Chen, Y., Li, Y., Tang, Y., Wang, F., and Xiang, W.: Local and~~
909 ~~distant source contributions to secondary organic aerosol in the Beijing urban area in~~
910 ~~summer, Atmos. Environ., 124, 176-185, <https://doi.org/10.1016/j.atmosenv.2015.08.098>,~~
911 ~~2016.~~
912 ~~Liu, F., Zhang, Q., Tong, D., Zheng, B., Li, M., Huo, H., and He, K. B.: High-resolution inventory of~~
913 ~~technologies, activities, and emissions of coal-fired power plants in China from 1990 to 2010, Atmos.~~
914 ~~Chem. Phys., 15, 13299-13317, 10.5194/acp-15-13299-2015, 2015J., Shen, J., Cheng, Z., Wang,~~
915 ~~P., Ying, Q., Zhao, Q., Zhang, Y., Zhao, Y., and Fu, Q.: Source apportionment and regional~~
916 ~~transport of anthropogenic secondary organic aerosol during winter pollution periods in the~~
917 ~~Yangtze River Delta, China, Sci. Total Environ., 710, 135620,~~
918 ~~<https://doi.org/10.1016/j.scitotenv.2019.135620>, 2020.~~
919 ~~Liu, X.-H., Zhang, Y., Cheng, S.-H., Xing, J., Zhang, Q., Streets, D. G., Jang, C., Wang,~~
920 ~~W.-X., and Hao, J.-M.: Understanding of regional air pollution over China using CMAQ,~~
921 ~~part I performance evaluation and seasonal variation, Atmos. Environ., 44, 2415-2426,~~
922 ~~<https://doi.org/10.1016/j.atmosenv.2010.03.035>, 2010.~~
923 ~~Liu, X., and Wang, J.: How important is organic aerosol hygroscopicity to aerosol indirect forcing?,~~
924 ~~Environ. Res. Lett., 5, 044010, 10.1088/1748-9326/5/4/044010, 2010.~~
925 ~~Liu, X. H., Zhu, Y. J., Zheng, M., Gao, H. W., and Yao, X. H.: Production and growth of new particles~~
926 ~~during two cruise campaigns in the marginal seas of China, Atmos. Chem. Phys., 14, 7941-7951,~~
927 ~~10.5194/acp-14-7941-2014, 2014.~~
928 ~~Luo, Y. X., Zheng, X. B., Zhao, T. L., and Chen, J.: A climatology of aerosol optical depth~~
929 ~~over China from recent 10 years of MODIS remote sensing data, Int. J. Climatol., 34, 863-~~
930 ~~870, 10.1002/joc.3728, 2014.~~
931 ~~Malm, W. C., Sisler, J. F., Huffman, D., Eldred, R. A., and Cahill, T. A.: Spatial and~~
932 ~~seasonal trends in particle concentration and optical extinction in the United States, J.~~
933 ~~Geophys. Res., 99, 1347-1370, doi:10.1029/93JD02916, 1994.~~
934 ~~Man, H., Zhu, Y., Ji, F., Yao, X., Lau, N. T., Li, Y., Lee, B. P., and Chan, C. K.: Comparison of Daytime and~~
935 ~~Nighttime New Particle Growth at the HKUST Supersite in Hong Kong, Environ. Sci. Technol., 49, 7170-~~
936 ~~7178, 10.1021/acs.est.5b02143, 2015.~~
937 ~~Massoli, P., Lambe, A. T., Ahern, A. T., Williams, L. R., Ehn, M., Mikkilä, J., Canagaratna,~~
938 ~~M. R., Brune, W. H., Onasch, T. B., Jayne, J. T., Petäjä, T., Kulmala, M., Laaksonen, A.,~~

939 Kolb, C. E., Davidovits, P., and Worsnop, D. R.: Relationship between aerosol oxidation
940 level and hygroscopic properties of laboratory generated secondary organic aerosol (SOA)
941 particles, *Geophys. Res. Lett.*, 37, 1-5, doi:10.1029/2010GL045258, 2010.

942 ~~Meyer, Murphy, B. N. K., Duplissy, J., Gysel, Woody, M., Metzger, A., Dommen, J., Weingartner, E.,~~
943 ~~Alfarra, M. R., Prevot, A. S. H., Fletcher, C., Good, N., McFiggans, Jimenez, J. L., Carlton, A. M.~~
944 ~~G., Jonsson, Hayes, P. L., Liu, S., Ng, N. L., Russell, L. M., Hallquist, M., Baltensperger,~~
945 ~~Setyan, A., Xu, L., Young, J., Zaveri, R. A., Zhang, Q., and Ristovski, Z. D.: Analysis of the~~
946 ~~hygroscopic Pye, H. O. T.: Semivolatile POA and volatile properties of ammonium sulphate~~
947 ~~seeded parameterized total combustion SOA in CMAQv5.2: impacts on source strength and~~
948 ~~unseeded SOA particles partitioning~~, *Atmos. Chem. Phys.*, 9, 721-732, doi:10.1117/1.11107-11133,
949 10.5194/acp-9-721-2009, 2009, 17-11107-2017, 2017.

950 Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan, R. C., and Seinfeld, J. H.:
951 Gas/Particle Partitioning and Secondary Organic Aerosol Yields, *Environ. Sci. Technol.*,
952 30, 2580-2585, doi:10.1021/es950943+, 1996.

953 Pankow, J. F.: An absorption model of gas/particle partitioning of organic compounds in
954 the atmosphere, *Atmos. Environ.*, 28, 185-188, [https://doi.org/10.1016/1352-](https://doi.org/10.1016/1352-2310(94)90093-0)
955 [2310\(94\)90093-0](https://doi.org/10.1016/1352-2310(94)90093-0), 1994.

956 Pankow, J. F., Marks, M. C., Barsanti, K. C., Mahmud, A., Asher, W. E., Li, J., Ying, Q.,
957 Jathar, S. H., and Kleeman, M. J.: Molecular view modeling of atmospheric organic
958 particulate matter: Incorporating molecular structure and co-condensation of water, *Atmos.*
959 *Environ.*, 122, 400-408, <https://doi.org/10.1016/j.atmosenv.2015.10.001>, 2015.

960 ~~Peter, T., Marcolli, C., Spichtinger, P., Corti, T., Baker, M. B., and Koop, T.: When Dry Air Is Too Humid,~~
961 ~~Science, 314, 1399-1402, doi:10.1126/science.1135199, 2006.~~

962 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopic
963 growth and cloud condensation nucleus activity, *Atmos. Chem. Phys.*, 7, 1961-1971,
964 10.5194/acp-7-1961-2007, 2007.

965 ~~Polichetti, G., Cecco, S., Spinali, A., Trimarco, V., and Nunziata, A.: Effects of particulate matter (PM10,~~
966 ~~PM2.5 and PM1) on the cardiovascular system, Toxicology, 261, 1-8,~~
967 ~~https://doi.org/10.1016/j.tox.2009.04.035, 2009.~~

968 ~~Poulain, L., Wu, Z., Petters, M. D., Wex, H., Hallbauer, E., Wehner, B., Massling, A., Kreidenweis, S. M.,~~
969 ~~and Stratmann, F.: Towards closing the gap between hygroscopic growth and CCN activation for~~
970 ~~secondary organic aerosols Part 3: Influence of the chemical composition on the hygroscopic~~
971 ~~properties and volatile fractions of aerosols, Atmos. Chem. Phys., 10, 3775-3785, doi:10.5194/acp-10-~~
972 ~~3775-2010, 2010.~~

973 Prisle, N. L., Engelhart, G. J., Bilde, M., and Donahue, N. M.: Humidity influence on gas-
974 particle phase partitioning of α -pinene + O₃ secondary organic aerosol, *Geophys. Res. Lett.*,
975 37, 1-5, doi:10.1029/2009gl041402, 2010.

976 Pye, H. O. T., Murphy, B. N., Xu, L., Ng, N. L., Carlton, A. G., Guo, H., Weber, R.,
977 Vasilakos, P., Appel, K. W., Budisulistiorini, S. H., Surratt, J. D., Nenes, A., Hu, W.,
978 Jimenez, J. L., Isaacman-VanWertz, G., Misztal, P. K., and Goldstein, A. H.: On the
979 implications of aerosol liquid water and phase separation for organic aerosol mass, *Atmos.*

980 Chem. Phys., 17, 343-369, 10.5194/acp-17-343-2017, 2017.

981 ~~Quaas, J., Boucher, O., Bellouin, N., and Kinne, S.: Satellite based estimate of the direct and indirect~~
982 ~~aerosol climate forcing, J. Geophys. Res., 113, D505204, doi:10.1029/2007JD008962, 2008.~~

983 Qiao, X., Ying, Q., Li, X., Zhang, H., Hu, J., Tang, Y., and Chen, X.: Source apportionment
984 of PM_{2.5} for 25 Chinese provincial capitals and municipalities using a source-oriented
985 Community Multiscale Air Quality model, Sci. Total Environ., 612, 462-471,
986 <https://doi.org/10.1016/j.scitotenv.2017.08.272>, 2018.

987 Ramanathan, V., Crutzen, P. J., Kiehl, J. T., and Rosenfeld, D.: Aerosols, Climate, and the
988 Hydrological Cycle, Science, 294, 2119-2124, 10.1126/science.1064034, 2001.

989 Rickards, A. M. J., Miles, R. E. H., Davies, J. F., Marshall, F. H., and Reid, J. P.:
990 Measurements of the Sensitivity of Aerosol Hygroscopicity and the κ Parameter to the O/C
991 Ratio, J. Phys. Chem. A, 117, 14120-14131, 10.1021/jp407991n, 2013.

992 ~~Rosenfeld, D., Lohmann, U., Raga, G. B., O'Dowd, C. D., Kulmala, R., Robinson, A. L., Donahue, N. M.,~~
993 ~~Fuzzi, S., Reissell, A., Shrivastava, M. K., Weitkamp, E. A., Sage, A. M., Grieshop, A. P., Lane,~~
994 ~~T. E., Pierce, J. R., and Andreae, M. O.: Flood or Drought: How Do~~
995 ~~Organic Aerosols Affect Precipitation?: Semivolatile Emissions and Photochemical Aging,~~
996 ~~Science, 321, 1309-1313, 1259-1262, 10.1126/science.1160606, 2008.~~
997 Pandis, S. N.: Rethinking
Science, 321, 1309-1313, 1259-1262, 10.1126/science.1160606, 2008.

997 Seinfeld, J. H., Erdakos, G. B., Asher, W. E., and Pankow, J. F.: Modeling the Formation
998 of Secondary Organic Aerosol (SOA). 2. The Predicted Effects of Relative Humidity on
999 Aerosol Formation in the α -Pinene-, β -Pinene-, Sabinene-, Δ^3 -Carene-, and Cyclohexene-
1000 Ozone Systems, Environ. Sci. Technol., 35, 1806-1817, 10.1021/es001765+, 2001.

1001 Shi, Z., Li, J., Huang, L., Wang, P., Wu, L., Ying, Q., Zhang, H., Lu, L., Liu, X., Liao, H.,
1002 and Hu, J.: Source apportionment of fine particulate matter in China in 2013 using a source-
1003 oriented chemical transport model, Sci. Total Environ., 601-602, 1476-1487,
1004 <https://doi.org/10.1016/j.scitotenv.2017.06.019>, 2017.

1005 Shrivastava, M., Cappa, C. D., Fan, J., Goldstein, A. H., Guenther, A. B., Jimenez, J. L.,
1006 Kuang, C., Laskin, A., Martin, S. T., Ng, N. L., Petaja, T., Pierce, J. R., Rasch, P. J., Roldin,
1007 P., Seinfeld, J. H., Shilling, J., Smith, J. N., Thornton, J. A., Volkamer, R., Wang, J.,
1008 Worsnop, D. R., Zaveri, R. A., Zelenyuk, A., and Zhang, Q.: Recent advances in
1009 understanding secondary organic aerosol: Implications for global climate forcing, Rev.
1010 Geophys., 55, 509-559, doi:10.1002/2016RG000540, 2017.

1011 Shrivastava, M. K., Lane, T. E., Donahue, N. M., Pandis, S. N., and Robinson, A. L.:
1012 Effects of gas particle partitioning and aging of primary emissions on urban and regional
1013 organic aerosol concentrations, 113, 10.1029/2007jd009735, 2008.

1014 Simon, H., and Bhave, P. V.: Simulating the Degree of Oxidation in Atmospheric Organic
1015 Particles, Environ. Sci. Technol., 46, 331-339, 10.1021/es202361w, 2012.

1016 ~~Slowik, J. G., Stroud, C., Bottenheim, J. W., Brickell, P. C., Chang, R. Y. W., Liggio, J., Makar, P. A., Martin,~~
1017 ~~R. V., Moran, M. D., Shantz, N. C., Sjostedt, S. J., van Donkelaar, A., Vlasenko, A., Wiebe, H. A., Xia, A.~~
1018 ~~G., Zhang, J., Leaitch, W. R., and Abbatt, J. P. D.: Characterization of a large biogenic secondary organic~~
1019 ~~aerosol event from eastern Canadian forests, Atmos. Chem. Phys., 10, 2825-2845, 10.5194/acp-10-~~
1020 ~~2825-2010, 2010.~~

1021 ~~Spracklen, D. V., Jimenez, J. L., Carslaw, K. S., Worsnop, D. R., Evans, M. J., Mann, G. W., Zhang, Q.,~~
1022 ~~Ganagaratna, M. R., Allan, J., Coe, H., McFiggans, G., Rap, A., and Forster, P.: Aerosol mass spectrometer~~
1023 ~~constraint on the global secondary organic aerosol budget, *Atmos. Chem. Phys.*, **11**, 12109–12136,~~
1024 ~~[10.5194/acp-11-12109-2011](https://doi.org/10.5194/acp-11-12109-2011), 2011.~~

1025 Sun, J., Liang, M., Shi, Z., Shen, F., Li, J., Huang, L., Ge, X., Chen, Q., Sun, Y., Zhang,
1026 Y., Chang, Y., Ji, D., Ying, Q., Zhang, H., Kota, S. H., and Hu, J.: Investigating the PM_{2.5}
1027 mass concentration growth processes during 2013–2016 in Beijing and Shanghai,
1028 *Chemosphere*, **221**, 452–463, <https://doi.org/10.1016/j.chemosphere.2018.12.200>, 2019.

1029 ~~Tritscher~~Sun, Y., Du, W., Fu, P., Wang, Q., Li, J., Ge, X., Zhang, Q., Zhu, C., Ren, L., Xu,
1030 W., Zhao, J., Han, T., Dommen, J., DeCarlo, P. F., Gysel, M., Barmet, P. B., Praplan, A. P., Weingartner,
1031 E., Prévôt, A. S. H., Riipinen, I., Donahue, N. M., and Baltensperger, U.: Volatility and hygroscopicity of
1032 aging~~Worsnop, D. R., and Wang, Z.: Primary and secondary organic aerosol aerosols in a~~
1033 ~~smog chamber~~Beijing in winter: sources, variations and processes, *Atmos. Chem. Phys.*, **11**,
1034 [11477–11496](https://doi.org/10.5194/acp-11-11477-2011)[16](https://doi.org/10.5194/acp-11-11477-2011), 8309–8329, [10.5194/acp-11-11477-2011](https://doi.org/10.5194/acp-11-11477-2011), [2011a](https://doi.org/10.5194/acp-11-11477-2011)[16](https://doi.org/10.5194/acp-11-11477-2011)–8309–2016, 2016.

1035 ~~Tritscher, T., Jurányi~~Sun, Y. L., Wang, Z., Martin, M., Chirico, R., Gysel, M., Heringa, M. F.,
1036 DeCarloFu, P. F., SierauQ., Yang, T., Jiang, Q., Dong, H. B., Prévôt, A. S. H., Weingartner, E., Li,
1037 J., and Baltensperger, U.: Changes of hygroscopicityJia, J. J.: Aerosol composition, sources and
1038 morphologyprocesses during ageing of diesel sootwintertime in Beijing, China, *Atmos. Chem.*
1039 *Phys.*, **13**, 4577–4592, [10.5194/acp-13-4577-2013](https://doi.org/10.5194/acp-13-4577-2013), 2013.

1040 Tkacik, D. S., Presto, A. A., Donahue, N. M., and Robinson, A. L.: Secondary Organic
1041 Aerosol Formation from Intermediate-Volatility Organic Compounds: Cyclic, Linear, and
1042 Branched Alkanes, *Environ. Res. Lett.*, **6**, 1–10Sci. Technol., **46**, 8773–8781, [10.1088/1748-](https://doi.org/10.1088/1748-9326/6/3/034026)
1043 [9326/6/3/034026](https://doi.org/10.1088/1748-9326/6/3/034026), 2011b[1021/es301112c](https://doi.org/10.1088/1748-9326/6/3/034026), 2012.

1044 Varutbangkul, V., Brechtel, F. J., Bahreini, R., Ng, N. L., Keywood, M. D., Kroll, J. H.,
1045 Flagan, R. C., Seinfeld, J. H., Lee, A., and Goldstein, A. H.: Hygroscopicity of secondary
1046 organic aerosols formed by oxidation of cycloalkenes, monoterpenes, sesquiterpenes, and
1047 related compounds, *Atmos. Chem. Phys.*, **6**, 2367–2388, [10.5194/acp-6-2367-2006](https://doi.org/10.5194/acp-6-2367-2006), 2006.

1048 Wang, K., Zhang, Y., Jang, C., Phillips, S., and Wang, B.: Modeling intercontinental air
1049 pollution transport over the trans-Pacific region in 2001 using the Community Multiscale
1050 Air Quality modeling system, *J. Geophys. Res.*, **114**, D04307, doi:10.1029/2008JD010807,
1051 2009.

1052 ~~Wang, S., Xing, J., Chatani, S., Hao, J., Klimont, Z., Cofala, J., and Amann, M.: Verification of~~
1053 ~~anthropogenic emissions of China by satellite and ground observations, *Atmos. Environ.*, **45**, 6347–~~
1054 ~~6358, <https://doi.org/10.1016/j.atmosenv.2011.08.054>, 2011.~~

1055 ~~Wang, X., Ye, X., Chen, H., Chen, J., Yang, X., and Gross, D. S.: Online hygroscopicity and chemical~~
1056 ~~measurement of urban aerosol in Shanghai, China, *Atmos. Environ.*, **95**, 318–326,~~
1057 ~~<https://doi.org/10.1016/j.atmosenv.2014.06.051>, 2014.~~

1058 Wiedensohler, A., Cheng, Y. F., Nowak, A., Wehner, B., Achtert, P., Berghof, M., Birmili,
1059 W., Wu, Z. J., Hu, M., Zhu, T., Takegawa, N., Kita, K., Kondo, Y., Lou, S. R.,
1060 Hofzumahaus, A., Holland, F., Wahner, A., Gunthe, S. S., Rose, D., Su, H., and Pöschl, U.:
1061 Rapid aerosol particle growth and increase of cloud condensation nucleus activity by

1062 secondary aerosol formation and condensation: A case study for regional air pollution in
1063 northeastern China, *J. Geophys. Res.*, 114, D00G08, doi:10.1029/2008JD010884, 2009.

1064 Wiedinmyer, C., Akagi, S. K., Yokelson, R. J., Emmons, L. K., Al-Saadi, J. A., Orlando,
1065 J. J., and Soja, A. J.: The Fire INventory from NCAR (FINN): a high resolution global
1066 model to estimate the emissions from open burning, *Geosci. Model Dev.*, 4, 625-641,
1067 10.5194/gmd-4-625-2011, 2011.

1068 Woody, M. C., Baker, K. R., Hayes, P. L., Jimenez, J. L., Koo, B., and Pye, H. O. T.:
1069 Understanding sources of organic aerosol during CalNex-2010 using the CMAQ-VBS,
1070 *Atmos. Chem. Phys.*, 16, 4081-4100, 10.5194/acp-16-4081-2016, 2016.

1071 ~~Xing, Y. F., Xu, Y. H., Shi, M. H., and Lian, Y. X.: The impact of PM_{2.5} on the human respiratory system,
1072 *J. Thorac. Dis.*, 8, E69-E74, 10.3978/j.issn.2072-1439.2016.01.10, 2016.~~

1073 ~~Yang, Y., Liao, H., and Lou, S.: Increase in winter haze over eastern China in recent decades: Roles of
1074 variations in meteorological parameters and anthropogenic emissions, *J. Geophys. Res.*, 121, 13,050-
1075 013,065, 10.1002/2016jd025136, 2016.~~

1076 Ying, Q., Cureño, I. V., Chen, G., Ali, S., Zhang, H., Malloy, M., Bravo, H. A., and Sosa,
1077 R.: Impacts of Stabilized Criegee Intermediates, surface uptake processes and higher
1078 aromatic secondary organic aerosol yields on predicted PM_{2.5} concentrations in the
1079 Mexico City Metropolitan Zone, *Atmos. Environ.*, 94, 438-447,
1080 <https://doi.org/10.1016/j.atmosenv.2014.05.056>, 2014.

1081 Ying, Q., Li, J., and Kota, S. H.: Significant Contributions of Isoprene to Summertime
1082 Secondary Organic Aerosol in Eastern United States, *Environ. Sci. Technol.*, **Environmental
1083 Science & Technology**, 49, 7834-7842, 10.1021/acs.est.5b02514, 2015.

1084 ~~Yue, D. L., Hu, M., Zhang, R. Y., Wu, Z. J., Su, H., Wang, Z. B., Peng, J. F., He, L. Y., Huang, X. F., Gong,
1085 Y. G., and Wiedensohler, A.: Potential contribution of new particle formation to cloud condensation
1086 nuclei in Beijing, *Atmos. Environ.*, 45, 6070-6077, <https://doi.org/10.1016/j.atmosenv.2011.07.037>,
1087 2011.~~

1088 Zhang, H., Hu, J., Kleeman, M., and Ying, Q.: Source apportionment of sulfate and nitrate
1089 particulate matter in the Eastern United States and effectiveness of emission control
1090 programs, *Sci. Total Environ.*, 490, 171-181,
1091 <https://doi.org/10.1016/j.scitotenv.2014.04.064>, 2014.

1092 Zhang, Q., Streets, D. G., Carmichael, G. R., He, K. B., Huo, H., Kannari, A., Klimont, Z.,
1093 Park, I. S., Reddy, S., Fu, J. S., Chen, D., Duan, L., Lei, Y., Wang, L. T., and Yao, Z. L.:
1094 Asian emissions in 2006 for the NASA INTEX-B mission, *Atmos. Chem. Phys.*, 9, 5131-
1095 5153, 10.5194/acp-9-5131-2009, 2009.

1096 ~~Zhang, X. Y., Sun, J. Y., Lin, W. L., Gong, S. L., Shen, X. J., and Yang, S.: Characterization of new particle
1097 and secondary aerosol formation during summertime in Beijing, China *Tellus B*, 63, 382-394,
1098 10.1111/j.1600-0889.2011.00533.x, 2011.~~

1099 Zhao, B., Wang, S., Donahue, N. M., Jathar, S. H., Huang, X., Wu, W., Hao, J., and
1100 Robinson, A. L.: Quantifying the effect of organic aerosol aging and intermediate-volatility
1101 emissions on regional-scale aerosol pollution in China, *Sci. Rep.*, 6, 28815,
1102 10.1038/srep28815, 2016a.

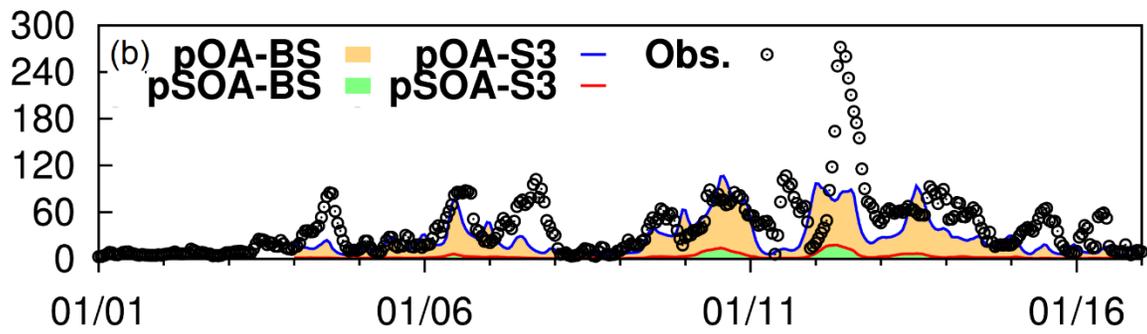
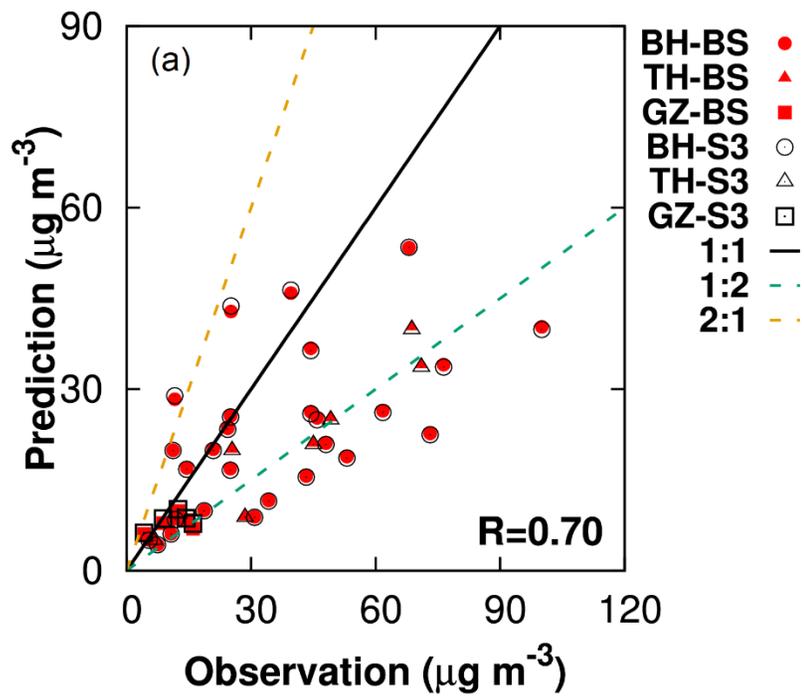
1103 Zhao, D. F., Buchholz, A., Kortner, B., Schlag, P., Rubach, F., Fuchs, H., Kiendler-Scharr,
1104 A., Tillmann, R., Wahner, A., Watne, Å. K., Hallquist, M., Flores, J. M., Rudich, Y.,
1105 Kristensen, K., Hansen, A. M. K., Glasius, M., Kourchev, I., Kalberer, M., and Mentel, T.
1106 F.: Cloud condensation nuclei activity, droplet growth kinetics, and hygroscopicity of
1107 biogenic and anthropogenic secondary organic aerosol (SOA), *Atmos. Chem. Phys.*, 16,
1108 1105-1121, 10.5194/acp-16-1105-2016, 2016b.

1109 ~~Zheng, B., Huo, H., Zhao, J., Qiu, Y., Zhou, W., Xu, W., Wang, J., Zhang, Q., Yao, Z., Li, L.,~~
1110 ~~Wang, Xie, C., Wang, Q., Du, W., Worsnop, D. R., Canagaratna, M. R., Zhou, L., Ge, X.-T.,~~
1111 ~~Yang, X. F., Liu, H., Fu, P., Li, J., Wang, Z., Donahue, N. M., and He, K. B.: High-resolution~~
1112 ~~mapping~~ Sun, Y.: Organic Aerosol Processing During Winter Severe Haze Episodes in
1113 Beijing, *J. Geophys. Res.*, 124, 10248-10263, 10.1029/2019jd030832, 2019.

1114 Zhao, Y., Hennigan, C. J., May, A. A., Tkacik, D. S., de Gouw, J. A., Gilman, J. B., Kuster,
1115 W. C., Borbon, A., and Robinson, A. L.: Intermediate-Volatility Organic Compounds: A
1116 Large Source of vehicle-emissions in Secondary Organic Aerosol, *Environ. Sci. Technol.*, 48,
1117 13743-13750, 10.1021/es5035188, 2014.

1118 Zheng, B., Zhang, Q., Zhang, Y., He, K. B., Wang, K., Zheng, G. J., Duan, F. K., Ma, Y.
1119 L., and Kimoto, T.: Heterogeneous chemistry: a mechanism missing in current models to
1120 explain secondary inorganic aerosol formation during the January 2013 haze episode in
1121 North China in 2008, *Atmos. Chem. Phys.*, 14, 9787-9805, 10.5194/acp-14-
1122 9787-2014, 2014 15-2031-2015, 2015.

1123



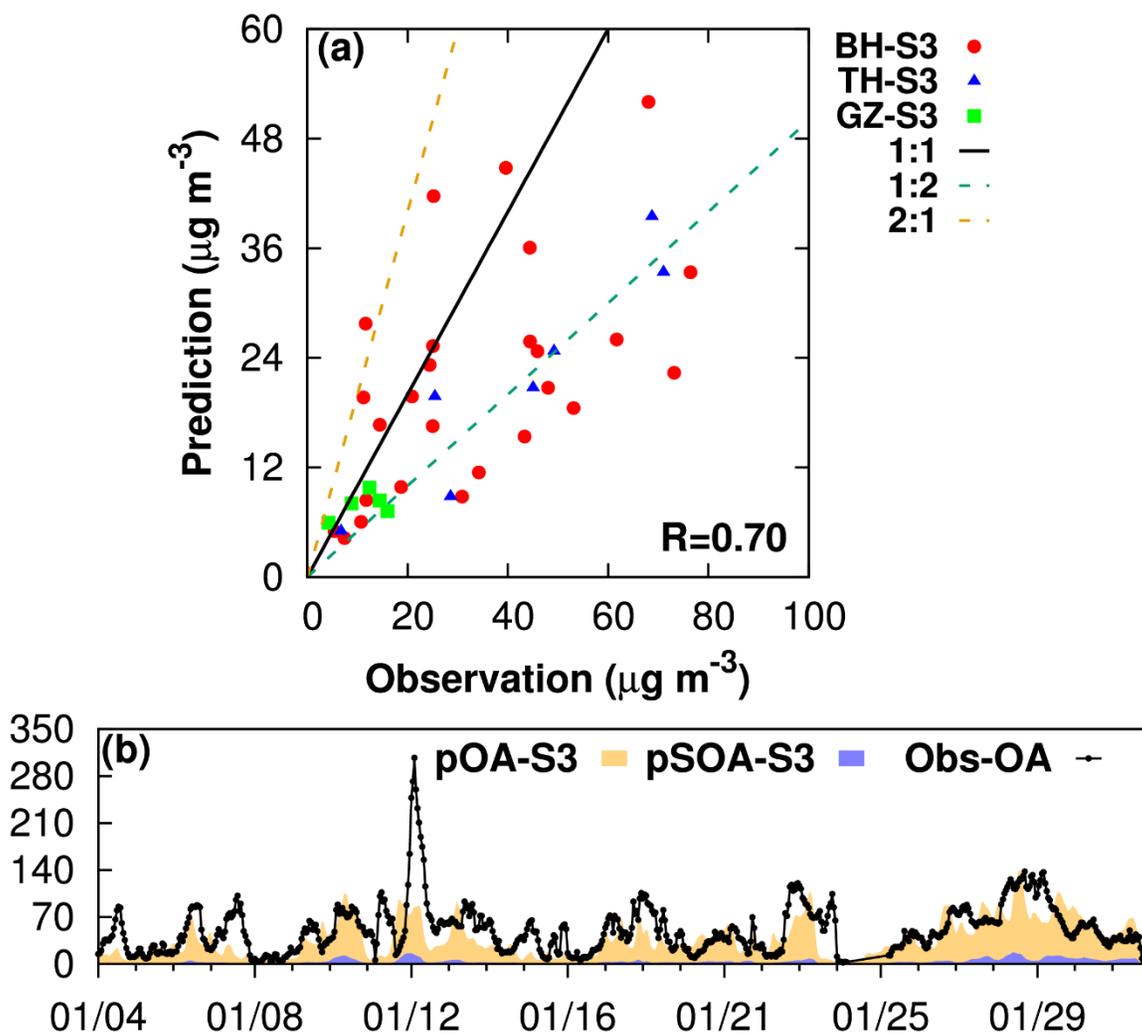


Figure 1. Comparison of (a) observed and modeled organic carbon concentration at University of Beihang (BH), Tsinghua University (TH) and Guangzhou (GZ); (b) observed organic aerosol (Obs-OA) at Beijing and predictions of total OA (pOA) and SOA (pSOA), unit is $\mu\text{g m}^{-3}$. Locations of monitoring sites are shown in Figure S1.

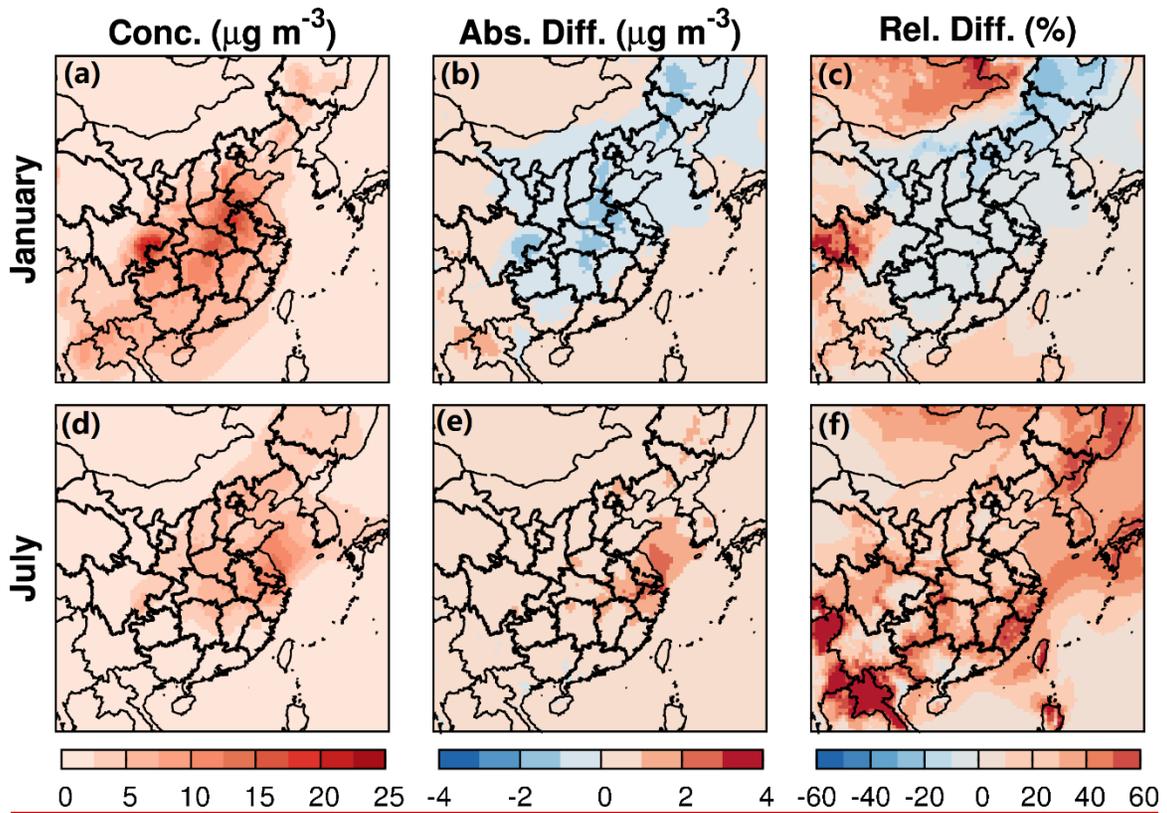
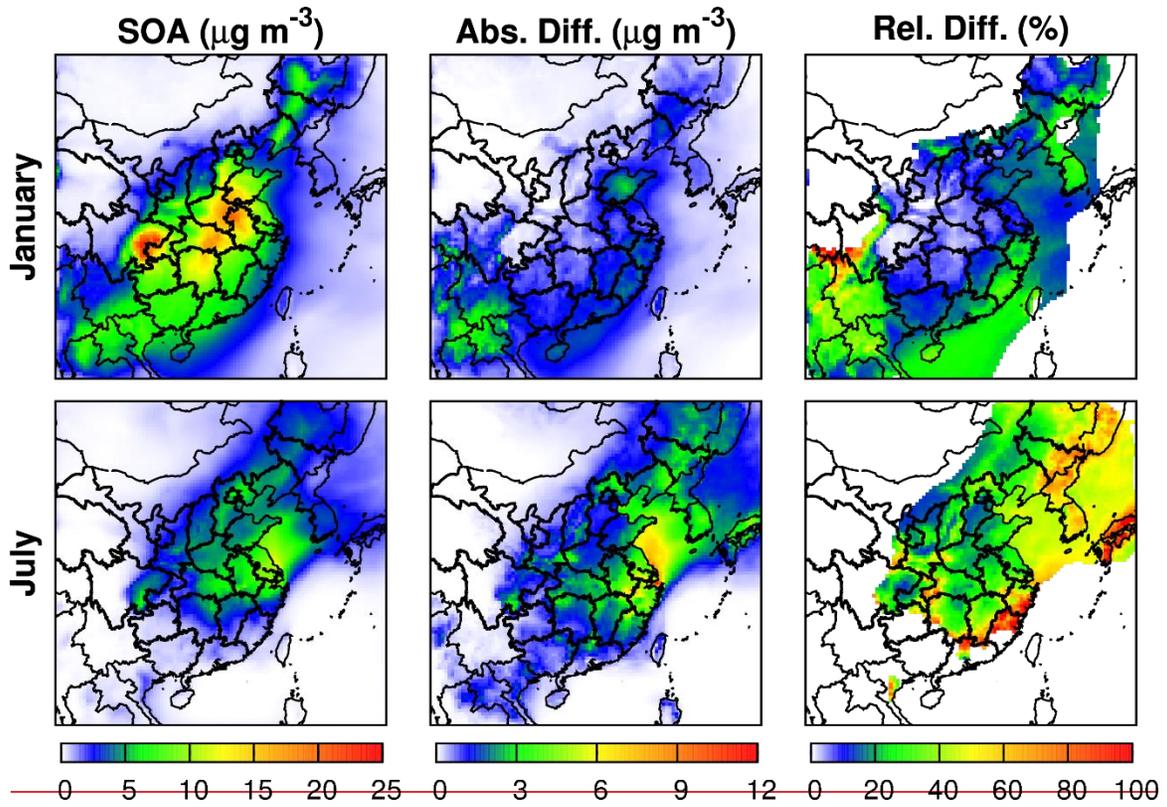
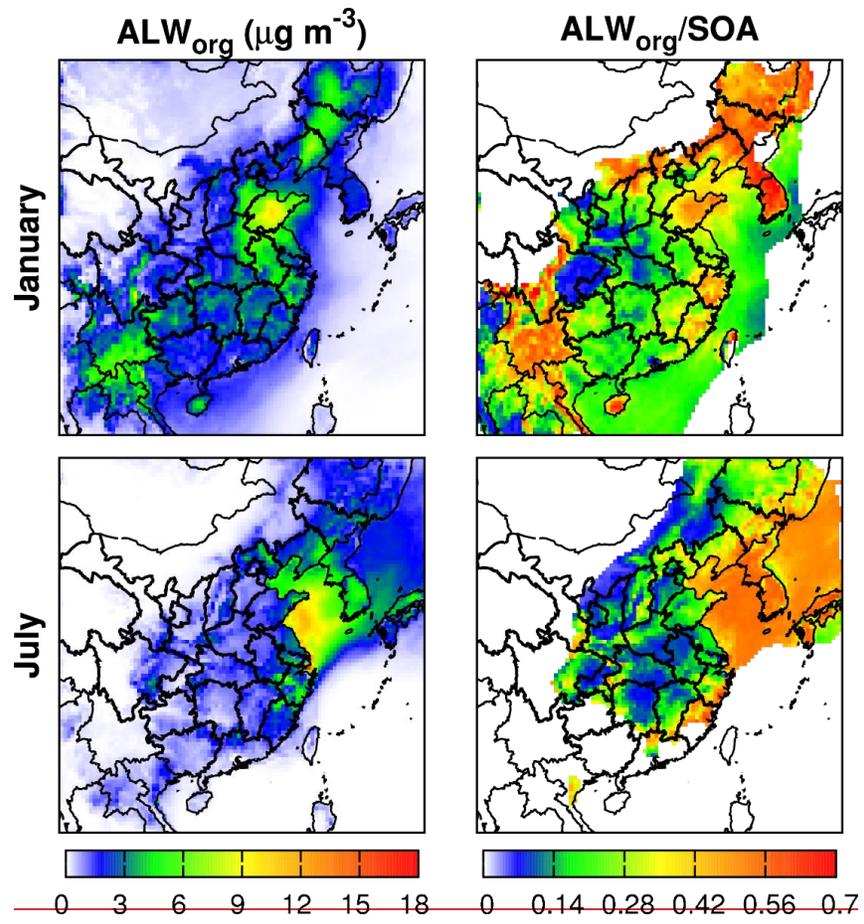


Figure 2. Monthly-averaged ~~total-SOA in BS~~ from S3 and ~~monthly-averaged-daily~~

~~maximum~~ changes of SOA due to water partitioning into OPM and non-ideality of the organic-water mixture. “Abs. Diff.” represents absolute differences (S3-BS); “Rel. Diff.” represents relative differences ((S3-BS)/BS, %). ~~Relative differences are shown in areas with monthly averaged SOA concentration greater than $1 \mu\text{g m}^{-3}$.~~



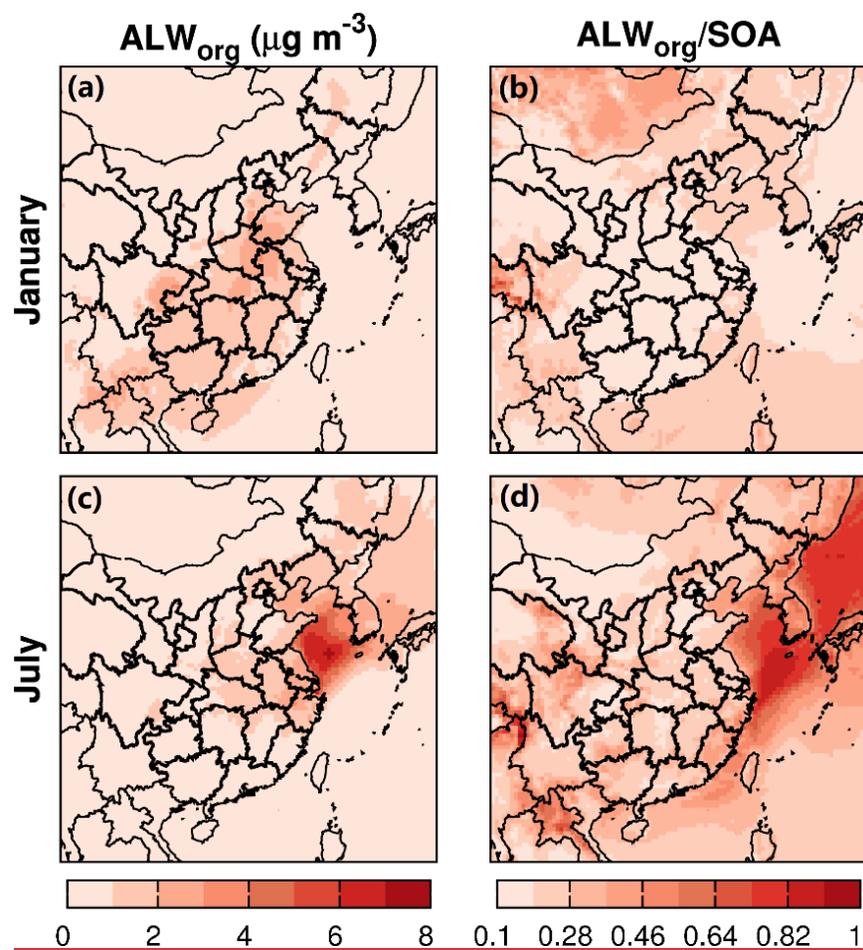
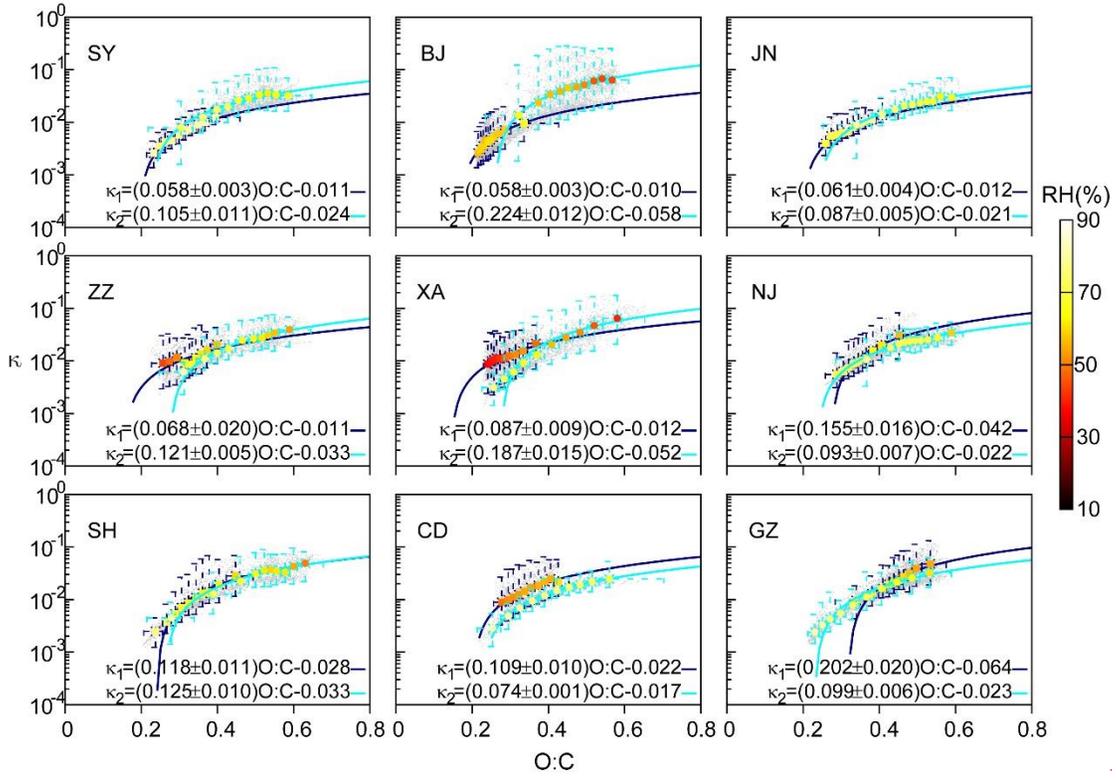


Figure 3. Monthly-averaged ~~daily maximum~~ water partitioning into the organic-phase (ALW_{org}, μg m⁻³) and the ratio to SOA (ALW_{org}/SOA) during January and July of 2013. ~~ALW_{org}/SOA is shown in areas with monthly-averaged SOA concentration greater than 1 μg m⁻³.~~



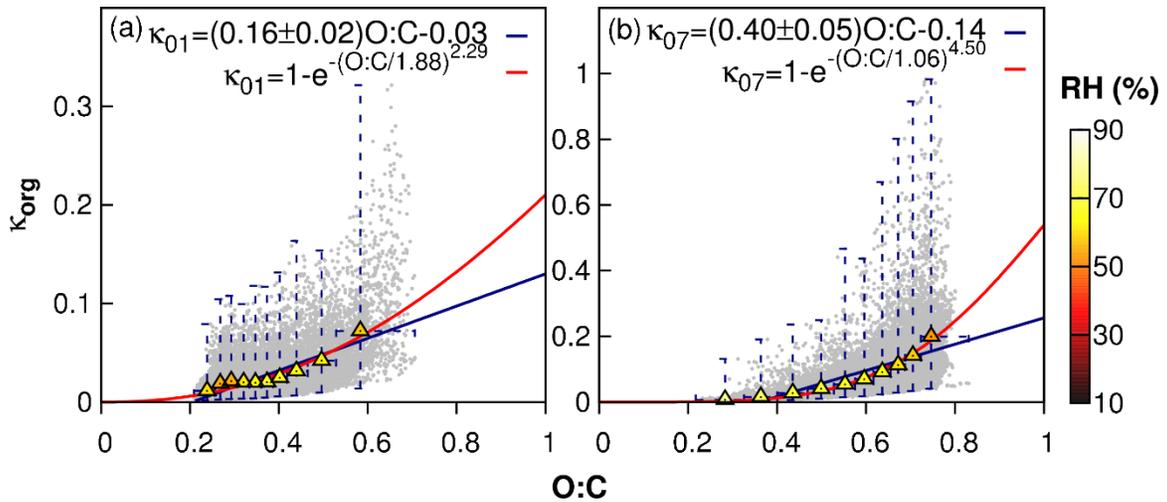


Figure 4. The correlation of hygroscopicity of organic aerosol (κ_{org}) and O:C ratio at in 9 representative cities including Shenyang (SS), Beijing (BJ), Jinan (JN), Zhengzhou (ZZ), Xi'an (XA), Nanjing (NJ), Shanghai (SH), Chengdu (CD), and Guangzhou (GZ). Gray dots on the background represent all the data in January (a) and July, which (b) of 2013. O:C ratios are categorized into several O:C10 bins. In each bin, the ranges of $\kappa_{O:C}$ and O:C ratio κ_{org} are represented by dashed bars colored for January (navy) and July (light blue), with the mean value. The mean values of O:C and κ_{org} are represented by triangles colored by the averaged RH of each bin. The mean κ The relationship between κ_{org} and O:C ratio are is fitted by a linear function with reduced major axis regression (blue lines) and an exponential function (red lines), respectively. κ_{01} and κ_{07} represent the fitted correlation for January and July, respectively.

| _____

|

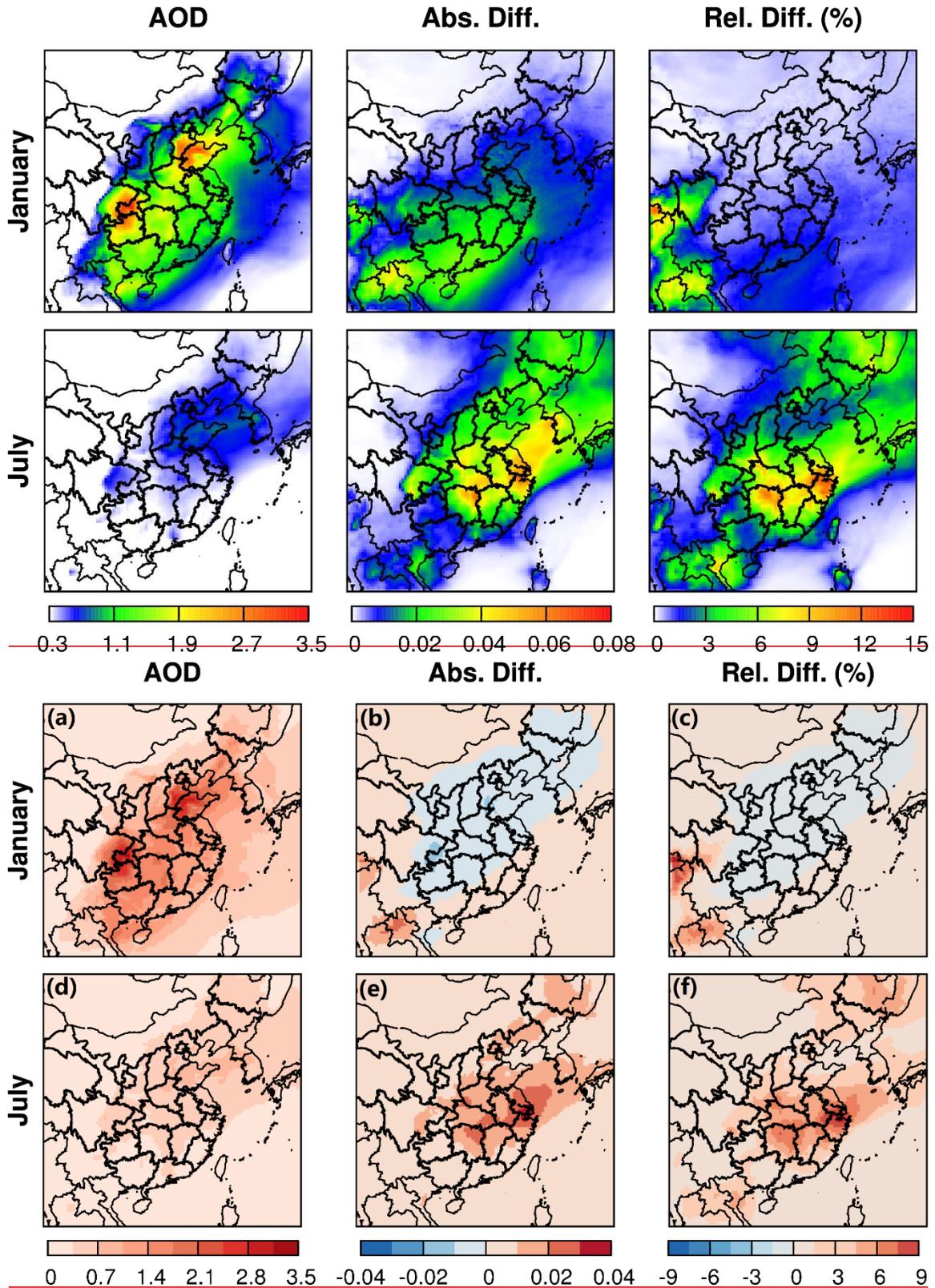
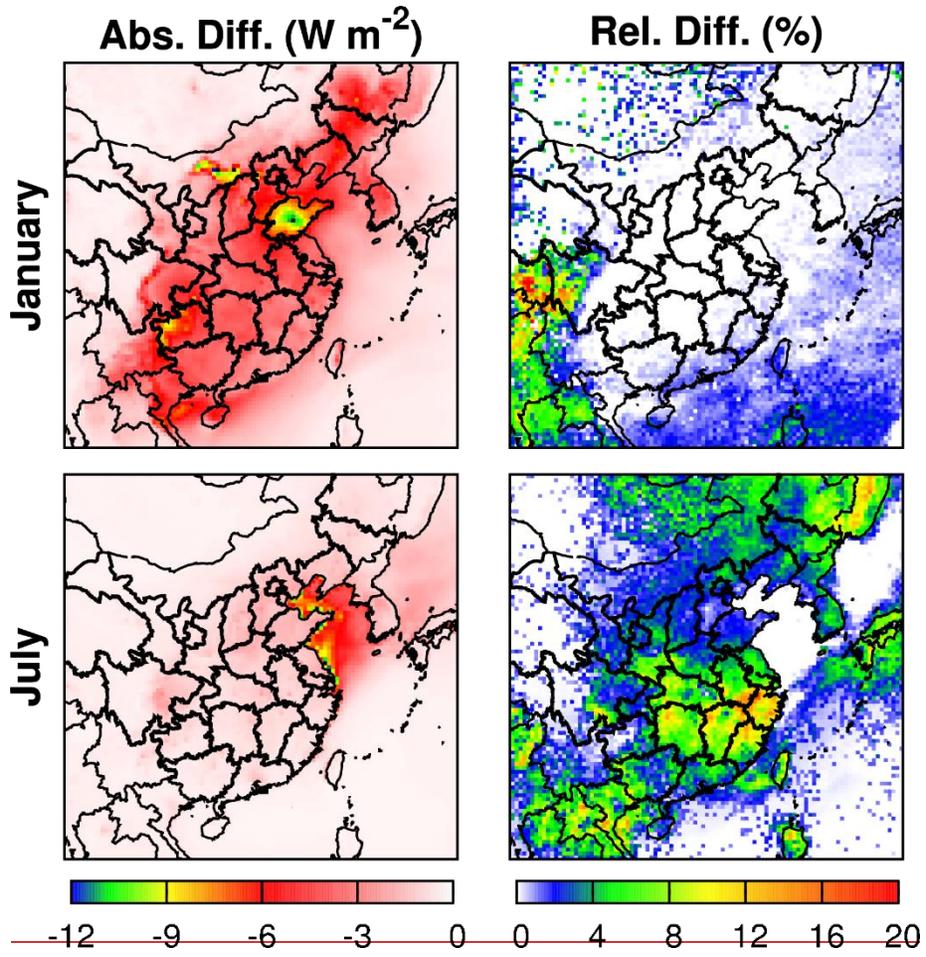


Figure 5. Monthly-averaged AOD_r at 550 nm and ~~the~~ monthly-averaged daily maximum changes of AOD_r due to water partitioning into OPM and ~~the~~ non-ideality of the organic-

water mixture. "Abs. Diff." represents absolute differences (S3-BS); "Rel. Diff." represents relative differences $((S3-BS)/BS, \%)$.



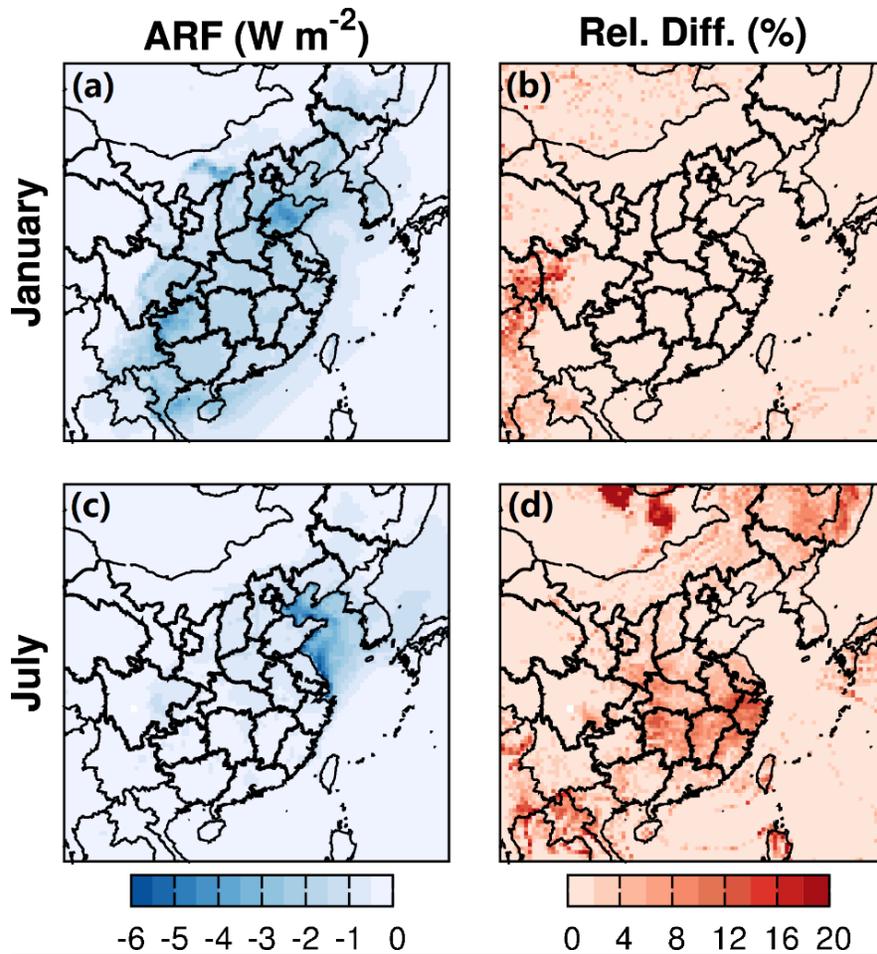
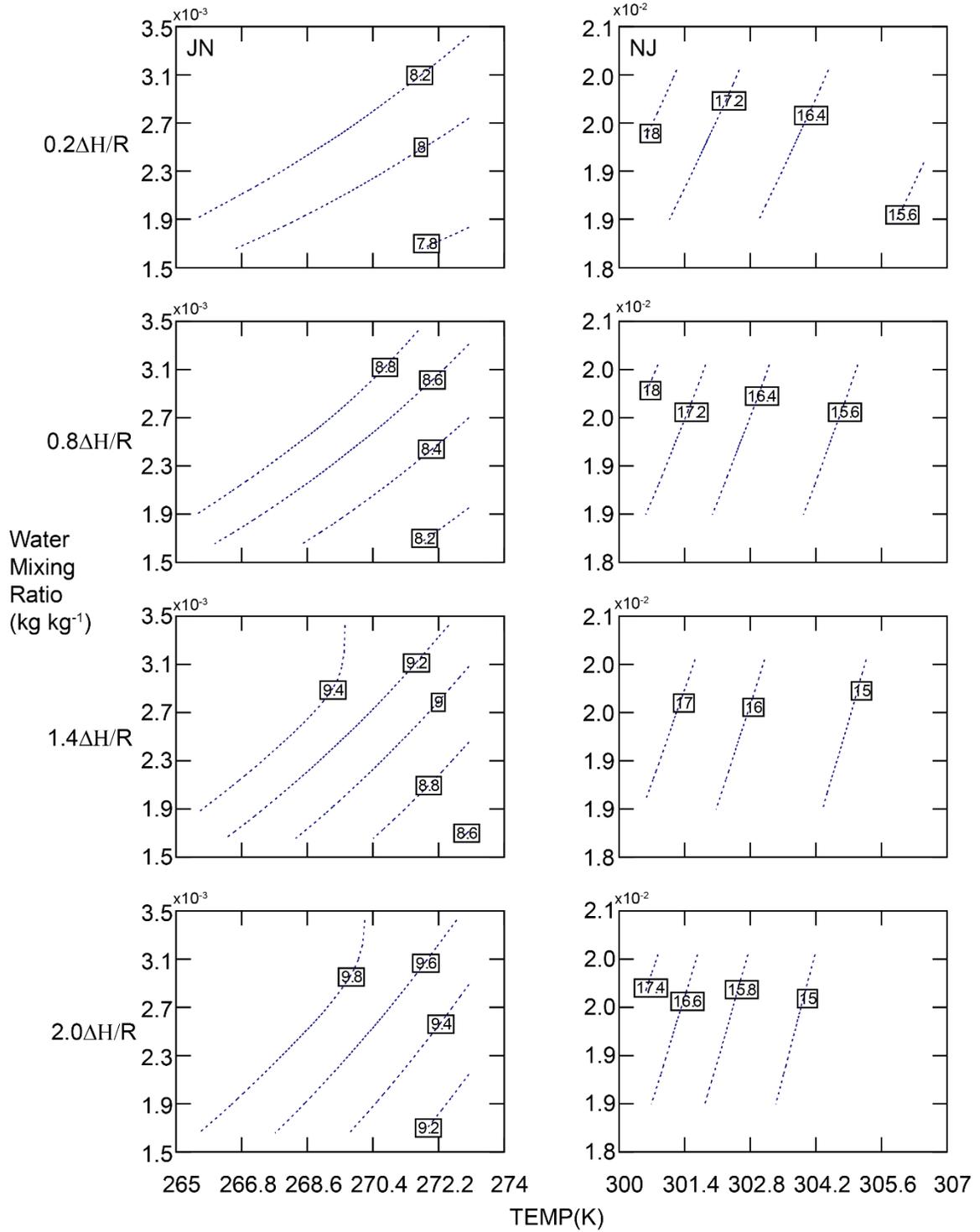


Figure 6. Monthly-averaged ~~daily maximum variation of~~ shortwave direct aerosol radiative forcing at the top of ~~the~~ atmosphere ~~from S3 and the relative changes~~ due to water partitioning ~~into OPM and non-ideality of the organic-water mixture~~ during January and July of 2013. ~~“Rel. Abs. Diff.” represents absolute differences (S3-BS); “Rel. Diff.”~~ represents relative differences ((S3-BS)/BS, %).



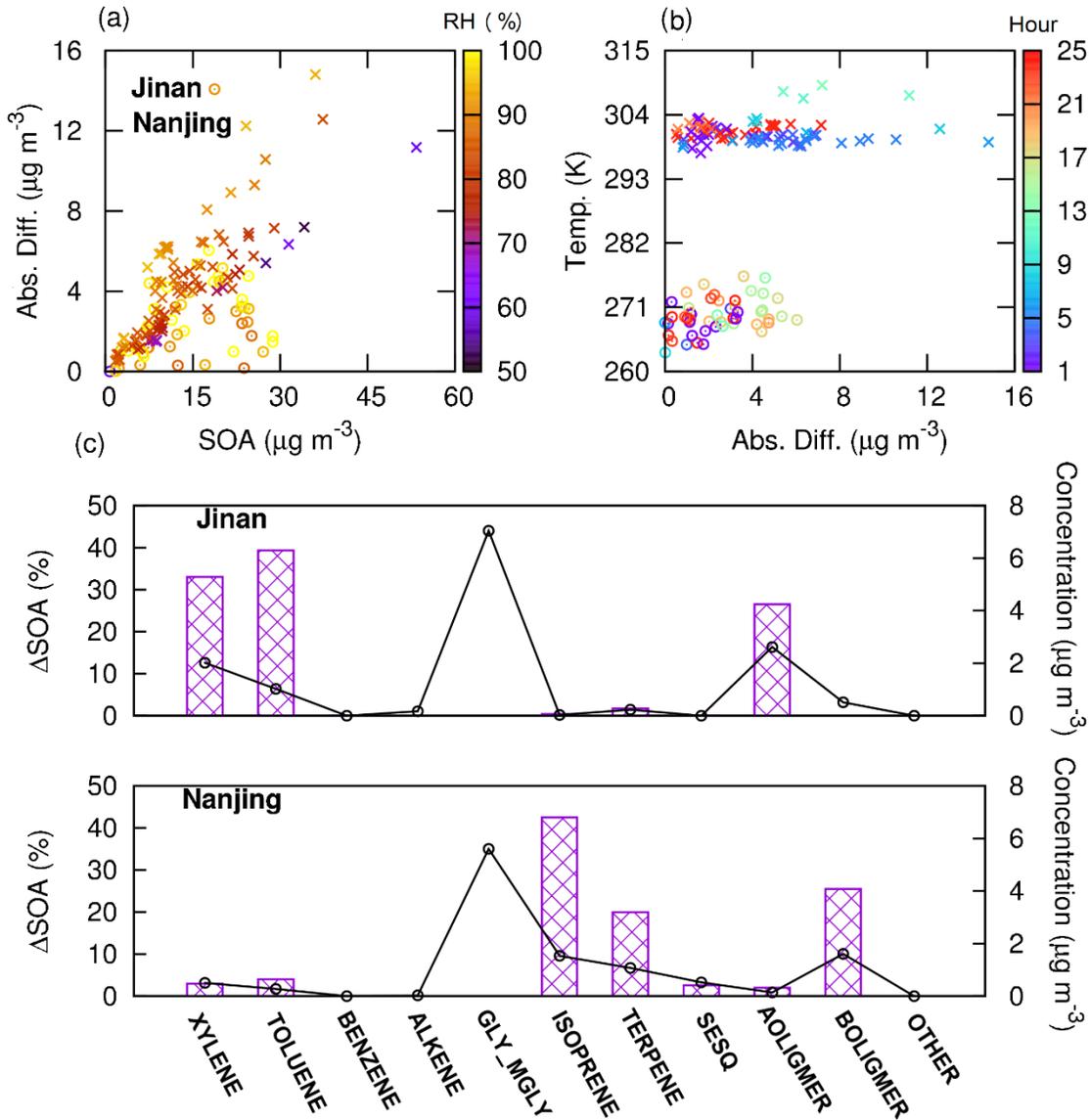


Figure 7. Correlation of water partitioning on SOA with (a) RH (b) temperature at Jinan in winter and Nanjing in summer, and the contribution from each SOA component to the total SOA increase. In plot (a) and (b), The sensitivity “Abs. Diff.” represents the daily maximum change of SOA that is calculated as $S_3 - B_S$. Color box represents RH in (a) and the hour in the day in (b) when daily maximum change of SOA occurred. In (c), the left axis represents contribution of each SOA component to the daily maximum SOA change due to water partitioning, and the right axis represents the concentration of each SOA component. —

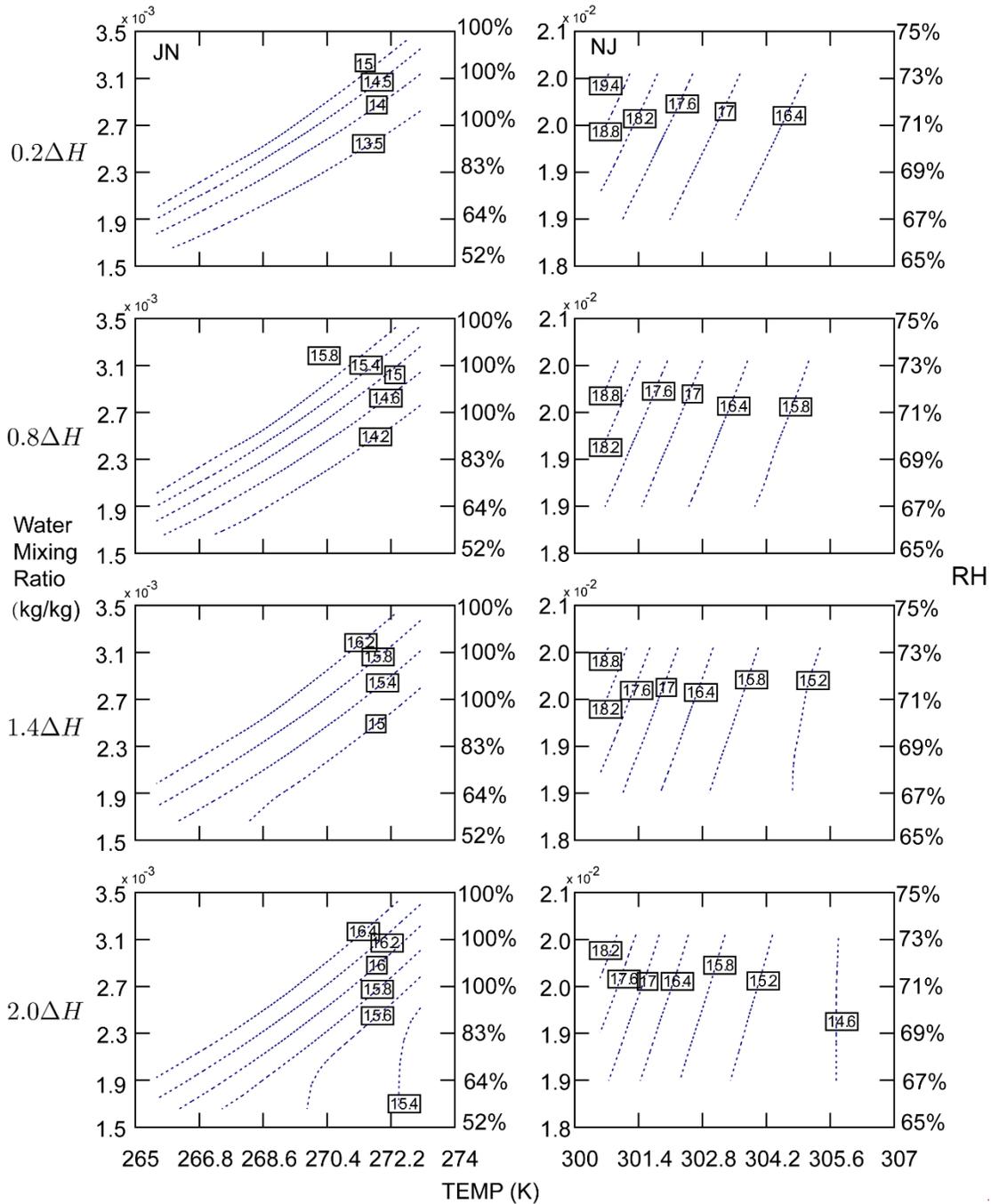


Figure 8. Sensitivity of SOA formation to temperature (TEMP), relative humidity (RH) water mixing ratio, and the temperature dependence parameter of SVP ($\Delta H/R$) at Jinan (JN, first column) and Nanjing (NJ, second column).

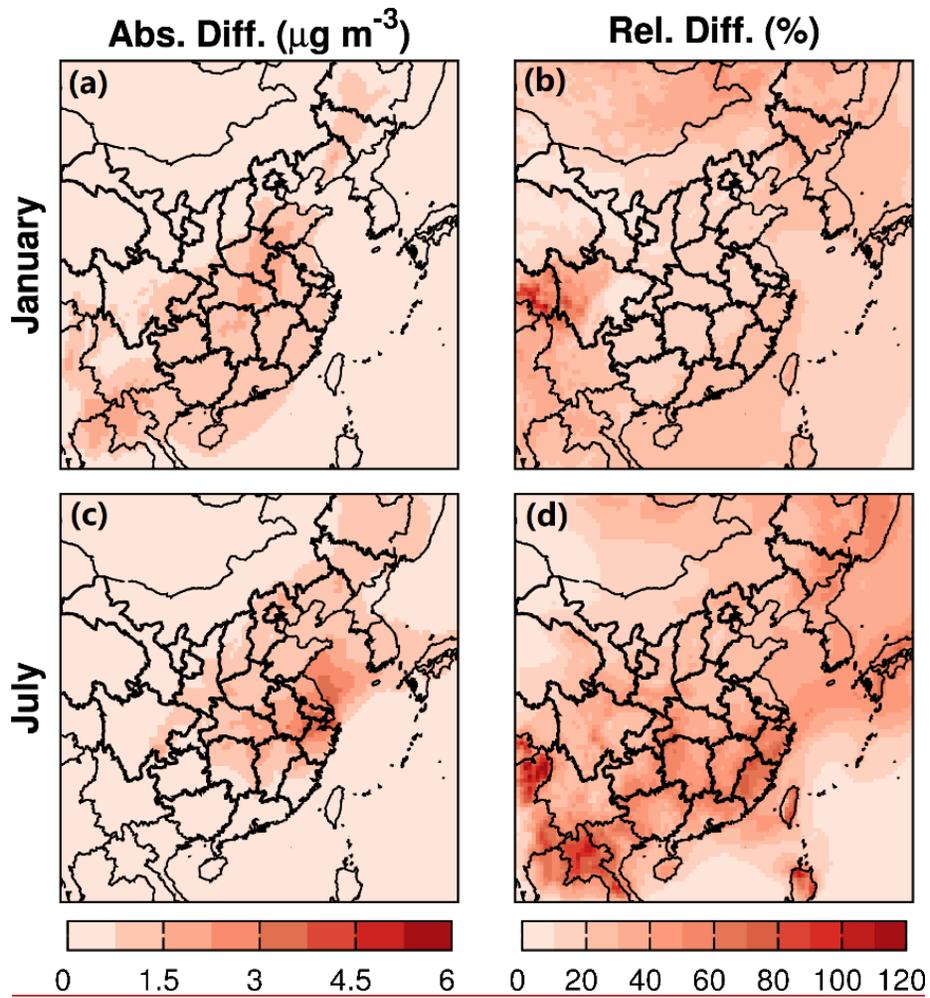


Figure 8. Monthly-averaged impacts of water partitioning into OPM on SOA. “Abs. Diff.” represents absolute differences ($S3-S2$); “Rel. Diff.” represents relative differences ($(S3-S2)/S2, \%$).

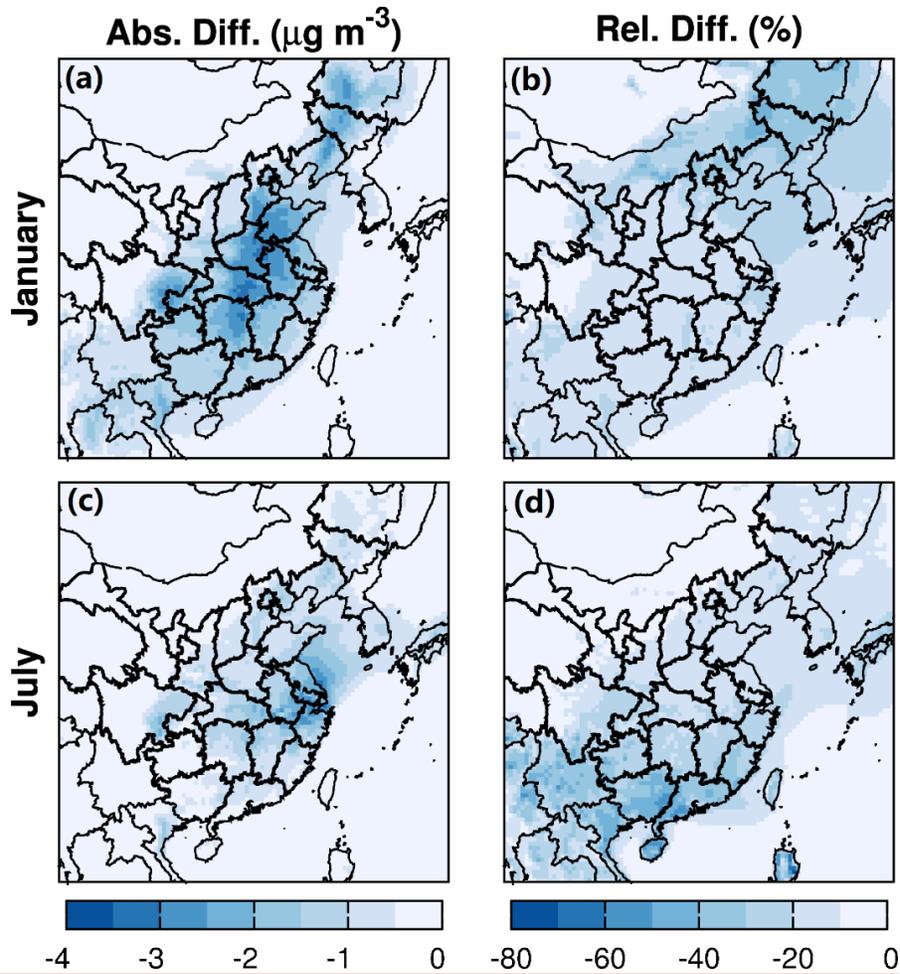


Figure 9. Monthly-averaged impacts of non-ideality of the organics-water mixture on SOA. “Abs. Diff.” represents absolute differences (S_3-S_1); “Rel. Diff.” represents relative differences $((S_3-S_1)/S_1, \%)$.