#### **Response to Referee #1**

The authors thank the reviewer for the comments that improve the quality of the paper. The detailed responses are given as follows. The reviewer comments are shown in italic font, the responses are in regular font, and the revised text is in **bold** font.

The authors present model results using the WRF/CMAQ for SOA formation in China. The model is updated for the partitioning of water vapor to the organic aerosol and the non-ideality of the organic phase. A comparison of the model to observations is performed for multiple sites. SOA enhancement during the summer and winter is discussed for the different China domains. The effect of aerosol liquid water on SOA formation and aerosol optical depth are presented. Correlation of the calculated OA hygroscopicity based on the k-Köhler theory to the O:C ratio is performed to show seasonal and multicity variations. Although the publication could provide valuable insights into the factors that govern SOA formation in China it currently lacks a detailed discussion and validation of the presented results. Therefore, the publication requires major revisions suggested below.

#### **Major Comments**

**Comment 1:** The manuscript is hard to read. Discussion and Results are not well separated in the manuscript. Many phrases are not clear and require more elaboration and better use of English. The table and most figures are poorly made and the science is hard to follow. Examples are given below.

**Response 1:** We thank the reviewer for pointing this out. We have updated all the figures and tables by showing monthly-averaged results instead of the daily maximum average as used in the previous version. Also, the text has been revised carefully to make it clear and easy to follow.

**Comment 2:** A major drawback in this work is that the model is not capturing the SOA formation during the winter that has been shown to be the dominant organic aerosol source in multiple publications for different domains of China. The authors only in a sentence discuss that the conversion of the POA to SOA may be the reason for these discrepancies but they have no observations to back this up.

**Response 2:** We agree with the reviewer that there is no direct observational evidence for this particular episode that POA aging played a significant role in the SOA formation. However, this process is one of the important missing

sources of SOA in several models, field and chamber studies (Robinson et al., 2007: Shrivastava et al., 2008: Zhao et al., 2016: Jimenez et al., 2009: Hodzic et al., 2010; Murphy et al., 2017). Also, organic compounds with intermediate volatility (IVOCs) between SVOCs and traditional VOCs, especially from combustion sources might contribute to SOA as well (Robinson et al., 2007; Shrivastava et al., 2008; Tkacik et al., 2012; Zhao et al., 2014; Zhao et al., 2016; Hodzic et al., 2010), but with high uncertainties in the emission inventory and SOA yields (Shrivastava et al., 2008; Tkacik et al., 2012). We performed an additional simulation using the latest version of the CMAQ model (v5.3.1), which includes a parameterization of these processes. The predicted SOA indeed increases significantly (Figure R1). We further analyzed the contribution of traditional POA (as semi-volatile POA in CMAQv5.3.1), SOA from the oxidation of the semi-volatile POA, traditional SOA, and a new SOA surrogate (pcSOA) representing missing SOA from IVOC oxidation, multigenerational aging of VOC oxidation products, and underestimate of SOA yield due to chamber wall losses (Murphy et al., 2017), finding pcSOA dominates in Beijing and Guangzhou (Figure R2) as well as the whole domain (Figure R1). The averaged SOA/POA ratio in Beijing is increased to 0.83, which is more consistent with field measurements (Zhao et al., 2019; Sun et al., 2016; Sun et al., 2013). However, the emission factors and oxidation rate of pcSOA precursors are highly uncertain and the contribution of pcSOA requires more observational constrains (Murphy et al., 2017).

We also examined the sensitivity of SOA and organic liquid water (ALW<sub>org</sub>) to pcSOA and POA in an offline calculation in Beijing, Guangzhou, Jinan, and Nanjing. POA has the same properties as we used in the model. Non-volatile isoprene SOA is taken to represent pcSOA as their similarities in saturation vapor pressure and O:C ratio. We found that both SOA and ALW<sub>org</sub> are positively correlated with pcSOA, increased by 2-5 times in different locations when pcSOA increased by 2 times. The impacts of water partitioning into OPM and non-ideality of the organic-water mixture by including the above process should be explored in a future study.

We extend the discussion of potential reasons for underestimated SOA in the revised text (L309-326):

"Again, no apparent changes of SOA nor OA are observed between case S3 and BS (not shown), since POA is predicted to be the primary contributor to OA at Beijing in winter in the current model, with an averaged SOA/POA ratio of 0.12. This ratio is much lower than the field observation of about 0.45-1.94 (Zhao et al., 2019; Sun et al., 2013; Sun et al., 2016). The bias might be due to the missing SOA converted by partitioning and aging of semi-volatile POA as well as oxidation from intermediate volatile organic compounds (IVOCs) and VOC oxidation products. Those pathways are shown to be important for SOA formation by modeling, field and chamber studies (Hodzic et al., 2010; Jimenez et al., 2009; Murphy et al., 2017; Robinson et al., 2007; Shrivastava et al., 2008; Tkacik et al., 2012; Zhao et al., 2014; Zhao et al., 2016a).

A sensitivity test was performed by using the newest CMAQ model version 5.3.1 that includes all the above processes in the aerosol module. The SOA/POA ratio in Beijing is improved greatly to be 0.83 in winter. However, high uncertainties still exist in the emissions of the involved precursors and characterization of SOA formation through these processes, needing further constrains by observations. Their influences on water partitioning into OPM and non-ideality of the organic-water mixture on SOA will be evaluated in a future study."



**Figure R1**. Mean SOA, SOA/OA, and pcSOA/SOA ratio predicted during January and July of 2013 by CMAQv5.3.1.



**Figure R2**. Modeled concentration of semi-volatile POA (sv-POA), OA and fraction of each organic aerosol component fsvPOA (sv-POA), foxPOA (oxidation of sv-POA), fpcSOA (pc-SOA) and fsvSOA (traditional SOA) in Beijing (a, c) and Guangzhou (b, d). Observations of OA in January 2013 at Beijing (Obs.) are also included in (a). The left axis is for the concentrations ( $\mu$  g m<sup>-3</sup>) and the right axis is the fraction of OA components.

**Comment 3:** There is no discussion on the influence of nitrate on aerosol liquid water. What fraction of the ALW is related to nitrate and what to organics? How could the ALW related to nitrate influence the partitioning of organics?

**Response 3:** In the current model, we separately treated the liquid water associated with organics (ALW<sub>org</sub>) and inorganics (ALW<sub>ing</sub>) in the condensed phase. Nitrate is assumed to only affect the inorganic aerosols and ALW<sub>ing</sub> as the interactions between inorganic and organic phases are not considered currently. This is the same approach used by Pankow et al. (2015). When considering the interactions between inorganic and organic aerosols in a CMAQ model, Pye et al. (2017) found an increase in SOA, which impacts are less significant than the separate treatment of the two phases. However, the interactions among condensed organics, i.e. the polarity of organics in the aerosol were ignored in their study. The interactions between inorganic and organic phases on ALW and SOA are beyond the scope of the current study and will be investigated in the future. We added a statement in the revised text (L207-L208) to make it clear:

# "In the current model, we assumed no interactions between the inorganic and organic phases."

**Comment 4:** A comparison of the model to observations should be performed and presented in the main text for both seasons in more detail. The effect of the improvements performed for the SAPRC-11 model is not discussed. The processes added e.g., the heterogeneous formation of nitrate and sulfate on the particle surface, SOA from isoprene, and dicarbonyls surface-controlled reactive uptake are not discussed. What is the effect of these added processes to the overall performance of the model? Each addition and the effect should be discussed in detail in order to support the importance of including them.

**Response 4:** The contribution of heterogeneous formation of nitrate and sulfate are not related to SOA formation. Papers documenting these changes have already been cited in the manuscript (Ying et al., 2014; Hu et al., 2016), and the impacts of the heterogeneous chemistry on nitrate and sulfate were discussed in those and another study (Zheng et al., 2015). The predicted nitrate and sulfate have been extensively compared with observations (Shi et al., 2017; Qiao et al., 2018). The improvements in the modeled SOA such as the reactive uptake of dicarbonyls and the isoprene generated epoxydiols have been discussed in previous studies by Ying et al. (2015), Li et al. (2015), Hu et al. (2017) and Liu et al. (2020) and have been shown to greatly increase the predicted SOA concentrations. The focus of this work is partitioning of water into OPM and the polarity of condensed organic compounds on SOA formation in China, which have not been examined so far. We have revised the manuscript to make it clear (L149-L161):

"Model configurations were largely based on that used by Hu et al. (2016) as summarized below. Firstly, SAPRC-11 was expanded for a more detailed treatment of isoprene oxidation and tracking dicarbonyl products (glyoxal and methylglyoxal) from different groups of major precursors (Ying et al., 2015). Secondly, SOA from isoprene epoxydiols (IEPOX), methacrylic acid epoxide (MAE) and dicarbonyls through surfacecontrolled irreversible reactive uptake were added (Hu et al., 2017; Li et al., 2015; Liu et al., 2020; Ying et al., 2015). Thirdly, the heterogeneous formation of secondary nitrate and sulfate from NO2 and SO2 reaction on the particle surfaces (Ying et al., 2014) were added, which is an important source of secondary inorganic aerosols (Zheng et al., 2015) and improves model estimates of nitrate and sulfate (Qiao et al., 2018; Shi et al., 2017). Fourthly, SOA yields were corrected for vapor wall loss (Zhang et al., 2014). Impacts of the above updates on model performances have been extensively discussed in the cited work and will not be further investigated in the current study."

**Comment 5:** *MEGAN has been shown to overestimate the isoprene emissions. Would this have a major effect on SOA formation in this work?*  **Response 5:** The MEGAN model has been shown to overestimate emissions of isoprene in the eastern and southeastern US (Wang et al., 2017; Kota et al., 2015), which is mainly due to databases of emission factors used. While our previous study indeed showed up to 5 times higher isoprene concentrations compared to the observations in Nanjing, the isoprene oxidation products (MACR and MVK) agree well with observations. Another study of Southern China also showed up to 2 times higher of MEGAN compared to measured isoprene fluxes (Situ et al., 2014). Thus, it is still inconclusive whether isoprene was indeed overestimated. If this isoprene overestimation was prevalent throughout the country and the isoprene SOA changed linearly with isoprene emissions, the actual SOA concentration and impacts on SOA due to water partitioning will decrease by 40-50% and 20-30%, respectively. Emissions of isoprene and other biogenic emissions are low during winter so that little or no impact is expected for winter.

**Comment 6:** Table 1: There is no discussion of the table in the main text. Abbreviations are not included in the caption. What is MB, GE, Num? Discussion on more statistically relevant values would be beneficial, e.g. what is the R2 of the comparisons? What are the presented values? Medians? Means? What is the domain of the model and how much do the values fluctuate around the domain? What are the uncertainties of the measurements and the model? A figure of the comparison of temperature and relative humidity of obs. vs model with the associated errors and linear regression analysis with the statistics would be informative.

**Response 6:** Table 1 shows the mean observation (OBS), mean prediction (PRE), mean bias (MB), gross error (GE), and the number of valid data of temperature and relative humidity in 8 sub-regions of the domain as shown in Figure S1. The table has been revised with explanations of all the abbreviations. A new table (Table S5) has been added to explain each region in Table 1. Since there are too many observation sites in the domain to show the uncertainties and regression information, we added R in the revised Table 1 and updated the table. More information about measurement methodology and uncertainties associated with measurements can be found at the NCDC data website. We discussed Table 1 in the original text (L245-L249):

"Table 1 shows the comparison of WRF predictions and observations in 8 subregions of the domain (Figure S1). Observed data are accessible from the National Climatic Data Center at <u>ftp://ftp.ncdc.noaa.gov/pub/data/noaa</u>. Temperature and RH are well captured by WRF in YRD, the Pearl River Delta (PRD), and central regions of China (the major regions of eastern China)."

We expanded this discussion in the revised manuscript in L277-L288:

"Temperature and RH are the two meteorological factors that affect SOA formation. Table 1 lists model statistics of mean observation (OBS), mean prediction (PRE), mean bias (MB), gross error (GE) and correlation coefficient (R) based on WRF and observations at monitoring sites located in 8 sub-regions of the domain (Figure S1) during January and July of 2013. The benchmarks for the MM5 model (another meteorology model) of 4-12km horizontal resolution suggested by Emery et al. (2001) are also listed in the table. Details of monitoring sites in the 8 sub-regions are listed in Table S5. Overall, WRF tends to underestimate both temperature and RH. The model shows better agreement with observed temperature as R is higher than that of RH. Both temperature and RH are well captured by the model in YRD, the Pearl River Delta (PRD), and the central regions of China (the major regions of eastern China). In these regions, MB and GE of temperature are -1.2~0.7 K and 1.8~2.6 K, respectively, which are -11.8~5.6% and 9.2~16.8% for RH, respectively."

**Comment 7:** Figure 1: (1) The current Figure has no information regarding the season the measurements are performed. A comparison of both seasons should be made and a figure like Fig1(a) should be made for each season. The time series should include the same site for the two seasons. (2) Why is the base case exactly the same as the S3 in Figure 1(b)? (3) Add errors to the measurements.

**Response 7:** (1) Due to limited observations of the simulated episode, we only have OC and OA measurements in January of 2013 at these sites. We used surface PM<sub>2.5</sub> alternatively to evaluate model performances in July of 2013. As a significant fraction of PM<sub>2.5</sub> in July is secondary, this still provides an indirect assessment of the model prediction of the oxidation capacity of the atmosphere, which is import for SOA formation. (2) The insignificant difference between BS and S3 in Figure 1(b) is likely due to a much smaller fraction of SOA compared to POA at this location predicted in the current model. Related discussions can refer to Response 2. (3) Unfortunately, there is no error information available for those measurements at this point.

**Comment 8:** Figure 2: Why do you use in (a) the base case and not the S3 case? I would consider promoting the updated S3 on the left and the changes on the middle and right panels.

**Response 8:** We have replaced case BS with S3 in Figure 2(a) and (d) and showed monthly-averaged results instead of averaged daily maximum in (b)-(c) and (e)-(f) in Figure 2.

**Comment 9:** Figure 3: I don't see the point in presenting the ratio of LWCorg to SOA. If both are expected to increase during pollution episodes then the ratio might stay the same therefore providing no valuable information. I would plot the SOA to OA as an alternative option or the SOA alone.

**Response 9:** One effect of water in the organic phase is that it decreases the average molecular weight of the absorbing organic phase, which could affect the subsequent partitioning of other semi-volatile organic compounds (see Eq 2). The ALW<sub>org</sub>/SOA allows the readers to see the importance to consider ALW<sub>org</sub> in the partitioning calculation as it can account for a significant fraction of SOA and lead to a reduced average molecular weight. We agree with the reviewer that this ratio might not be very different between clean and polluted episodes. However, it is nonetheless useful in assessing the importance of including ALW<sub>org</sub> in SOA modeling. SOA/OA is certainly very useful information as well. Since POA is still significantly higher than SOA, especially during the winter month, SOA/OA ratio did not change significantly when ALW<sub>org</sub> is considered as shown in Figure R3. We have added this figure in the revised supplemental materials as Figure S13.



**Figure R3**. Averaged SOA/OA ratio from case BS and S3 during January and July of 2013.



#### colors and increase font size.

**Response 10:** Figure 4 (as shown below) has been revised to show all the data from each city in each month. Detailed results of each city are shown in Figure S9 and S10.



**Figure 4.** The correlation of hygroscopicity of organic aerosol ( $\kappa_{org}$ ) and O:C

ratio at 9 representative cities including Shenyang (SS), Beijing (BJ), Jinan (JN), Zhengzhou (ZZ), Xi'an (XA), Nanjing (NJ), Shanghai (SH), Chengdu (CD), and Guangzhou (GZ) in January (a) and July (b) of 2013. O:C ratios are categorized into 10 bins. In each bin, the ranges of O:C and  $\kappa_{org}$  are represented by bars.

The mean values of O:C and  $\kappa_{org}$  are represented by triangles colored by the

averaged RH of each bin. The relationship between  $\kappa_{org}$  and O:C is fitted by

a linear function with reduced major axis regression (blue lines) and an exponential function (red lines), respectively.  $\kappa_{01}$  and  $\kappa_{07}$  represent the fitted correlation for January and July, respectively.

**Comment 11:** Figure 7: The graph is not clear. In the main text, the authors discuss that the daily maximum of SOA occurs when RH is greater than 70% in both cities. The RH is higher than 70% all the time. The time of the day is up to 24 hours and not 25. The markers and boxes for (c) are not discussed whether they represent the left or right axis.

**Response 11:** We have removed this figure to avoid confusion.

#### **Specific Comments**

Comment 12: Line 34: Please elaborate more

Response 12: Now the text reads:

"However, the models typically assume that the organic particulate matter (OPM) is an ideal mixture and ignore the partitioning of water vapor to OPM."

**Comment 13:** *Line 39: Please define generally with a statistical value that has meaning.* 

Response 13: The text has been revised in L40-42 as follows:

"The modified model can generally capture the observed surface organic carbon (OC) with a correlation coefficient R of 0.7, and the surface OA with the mean fractional bias (MFB) and mean fractional error (MFE) of - 0.28 and 0.54, respectively."

**Comment 14:** *Line 91: Please elaborate more on the "purer condensed organics" for non-experts.* 

**Response 14:** This should be condensed organics and has been revised accordingly.

**Comment 15:** *Line 99: "neglect" instead of "neglected".* **Response 15:** The text has been revised as instructed.

Comment 16: Line 99: Please elaborate more on 1). Response 16: The text has been revised in L98-100 as following: "1) the molecular structures and interactions of functional groups (-OH, -C=O, -COOH, etc.) of condensed organics (non-ideality);"

**Comment 17:** *Line 109: The sentence is missing a verb.* **Response 17:** The sentence has been revised in L109-111 as following:

"Laboratory and field studies have observed water absorbed by SOA from a variety of precursor VOCs (Lambe et al., 2011; Zhao et al., 2016b; Asa-Awuku et al., 2010; Varutbangkul et al., 2006)."

Comment 18: Line 128-131: The sentence is hard to read. Please rephrase. Response 18: The sentence has been revised in L126-129 as follows: "Using UNIversal Functional Activity Coefficient (UNIFAC) method (Fredenslund et al., 1975) for calculating activity coefficients of the organic-water mixture, it was found that in the eastern U.S., where biogenic SOA dominated the OA, considering ALW<sub>org</sub> leads to a significant increase in predicted SOA (Pankow et al., 2015; Jathar et al., 2016)."

#### Comment 19: Line 138: Which region?

**Response 19:** The region refers to China. The text has been revised in L134-135 as follows:

"Previous modeling studies in China indicate that SOA was underpredicted (Lin et al., 2016; Jiang et al., 2012)"

**Comment 20:** Line 163: (1) Don't change a line. (2) Also, why acidic conditions? Please, elaborate more. (3) Why is the reactive uptake of dicarbonyls, IEPOX and MAE in the "non-volatile" category?

**Response 20:** (1) This is a mistake due to file format conversion and has been corrected. (2) This is a typo. SOA formed by isoprene oxidation under acidic conditions refers to IEPOX and MAE SOA based on chamber experiments (Lal et al., 2012; Lin et al., 2012; 2013). We have removed "isoprene oxidation under acidic conditions" in the revised text. (3) In the current model, we assume that the reactive uptake of dicarbonyls, IEPOX and MAE is irreversible, as an upper-limit estimation of SOA from these precursors. That's why they are classified as non-volatile SOA. The text has been revised in L165-173 as following:

"SOA from dicarbonyls, IEPOX, and MAE were formed by irreversible reactive uptake and categorized as NV-SOA in the current model as well. Some studies investigated SOA from glyoxal, methylglyoxal, and IEPOX using detailed reactions and reversible pathways in models or observed as reversible processes in chamber experiments, leading to a relatively lower SOA yield compared to the surface-controlled irreversible uptake (Lim et al., 2013; Knote et al., 2014; Galloway et al., 2009; EI-Sayed et al., 2018; Budisulistiorini et al., 2017). The non-volatile assumption used in this paper allows an upper-limit estimation of the importance of these additional SOA formation pathways."

#### Comment 21: Line 164: change to "was mostly"

**Response 21:** This sentence has been moved to L162 in the revised manuscript and now it reads as follows:

"The SOA module mostly follows Pankow et al. (2015)."

**Comment 22:** *Line 165: Please elaborate more on the non-ideality calculation of the organic-water mixture for non-experts.* 

**Response 22:** This sentence has been moved to L199-200 in the revised manuscript and now it reads as follows:

"POA is also involved in the calculation of activity coefficients for the

organic-water mixture."

**Comment 23:** *Line 170: Is this the absorbing organic phase?* **Response 23:** Yes. The text has been revised to "the absorbing organic phase".

**Comment 24:** *Line 196: Change to "as water condenses".* **Response 24:** The text has been revised as instructed.

**Comment 25:** Line 200: Please elaborate more on the "Kelvin effect neglected" for non-experts.

**Response 25:** The text has been revised in L210-211 as following:

"Based on the κ-Köhler theory with linearly additive hygroscopic behavior of each component of the mixed particle"

**Comment 26:** *Line 204: Change to "can be estimated".* **Response 26:** The text has been revised as instructed.

**Comment 27:** Line 242: Observations in 8 sub-regions of the domain during which period?

**Response 27:** The text has been revised in L278-281 as follows:

"Table 1 lists model statistics of mean observation (OBS), mean prediction (PRE), mean bias (MB), gross error (GE) and correlation coefficient (R) based on WRF and observations at monitoring sites located in 8 sub-regions of the domain (Figure S1) during January and July of 2013."

**Comment 28:** Line 251: No significant improvements observed when applying the above additions means that the model is still missing a significant pathway to SOA formation, especially since OA in both seasons are dominated by SOA based on observations. This should be discussed in detail and in the context of previous studies and findings from AMS measurements in China.

**Response 28:** We have revised the text and added a discussion about the underestimation of SOA in the current model. Please refer to Response 2 for more details.

**Comment 29:** Line 258: Does it capture the observed diurnal variation? What is the R2 or R of the timeseries of the modeled to observed values? What is the ratio of the two? In many cases, it seems that the difference is higher than a factor of 4. Is that a usual discrepancy? If so, how is much is it improved when incorporating the detailed SOA models?

Response 29: Our model can capture the diurnal variation. The mismatching

of several peak values might be due to uncertainties in the emission inventory and the underestimate of SOA in the current model. The R of the modeled to observed OA is 0.55. The ratio of the averaged prediction to observation is 0.75. Since the model predicts a very small ratio of SOA to POA, the improvement from the detailed SOA model is insignificant. The small SOA/POA ratio might be due to the missing SOA from other pathways including POA aging and oxidation from IVOCs and VOC oxidation products. This has been explained in Response 2.

We revised the text in L301-304 to expand more discussion of the modeled and observed OA comparison:

"CMAQ can well capture the observed diurnal variation of OA in Beijing during wintertime, except for the underestimates of peak values. The correlation coefficient of modeled to observed OA is 0.55. We find a 25% underestimate of OA on average."

**Comment 30:** Line 259: It would be great if "better" was described with statistical terminology. A way to describe the data and the comparison to modeled values would be to generate box and whiskers of the ratio of observations to modeled values for non-polluted and polluted days, respectively.

**Response 30:** We did a mistake in the MFB and MFE calculation for OA, which should be -0.28 and 0.54, respectively. We also calculated the biases on polluted and non-polluted days of OA. MFB and MFE of polluted days are -0.38 and 0.64, which are -0.26 and 0.52 for non-polluted days. The text has been revised in L306-309 as follows:

"The mean fractional bias (MFB) and mean fractional error (MFE) of polluted days are -0.38 and 0.64, respectively, which are worse than that of the non-polluted days (-0.26 for MFB and 0.52 for MFE). The overall MFB and MFE of OA during January are -0.28 and 0.54, within the criteria (MFB $\leq$ ±0.6; MFE $\leq$ 0.75) suggested by EPA (2007)."

**Comment 31:** Line 262-263: POA is not the primary contributor to OA in Beijing in winter. Many studies show that SOA is the major contributor and the path towards SOA formation is currently unknown and strongly dependent on LWC in the particles. Aging of POA not treated in the model is not guaranteed to be the main source of SOA.

**Response 31**: Please refer to Response 2.

**Comment 32:** Line 280: Here only one season is provided in terms of timeseries comparison of the model and obs. Please provide both seasons. **Response 32:** Unfortunately, detailed chemical composition measurements for

aerosols are very limited in China during 2013. We only have observations of OC and OA in January of 2013 and  $PM_{2.5}$  in July of 2013 available for model evaluation.

**Comment 33:** Line 285: Figure S5 shows the anthropogenic SOA and not the dicarbonyl SOA. Please separate the contributions and discuss them in the main text. Identifying the contribution of different compounds to SOA formation in China would be of great interest to the scientific community.

**Response 33:** The contribution of each precursor to SOA of this episode has been shown in Hu et al. (2017) and will not be discussed in detail in the current study. The text has been revised in L344-346 as following:

"Anthropogenic emissions are the major sources of SOA (Figure S6), such as dicarbonyl products from the oxidation of xylene and toluene (Hu et al., 2017).

#### Comment 34: Line 312: What about particulate nitrate?

**Response 34:** The interactions between water-inorganics and water-organics are treated separately in the current model. We only focus on the water-organic interaction in the current study.

**Comment 35:** Line 380: RH is higher than 70% all the time. What is the meaning of this sentence?

**Response 35:** We have removed this figure to avoid confusion.

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#### **Response to Referee #2**

The authors thank the reviewer for the comments that improve the quality of the paper. The detailed responses are given as follows. The reviewer comments are shown in italic fonts, the responses are in regular font, and the revised text is in **bold** font.

In this paper, Li et al. have modified the CMAQ model to take into account the impacts of water partitioning and polarity of organic compounds on SOA formation. The model was applied over Eastern China to estimate the regional and seasonal impacts of these modifications on SOA and the aerosol water content. This study may have potential to contribute in the organic aerosol modeling field but major revisions needs to be done before publication. In particular, I have several concerns regarding the validity of the scientific methodology used and the presentation of the study. Therefore, I would recommend publication only if these comments will be addressed and fundamental changes will be contacted.

#### **Major comments**

**Comment 1:** Page 4 line 94: The majority of current CTMs have replaced the 2-product model with the VBS approach. Please make this clear and refer to the 2-product model of Odum et al. (1996) for historical reasons.

**Response 1:** The text has been revised in the manuscript (L88-95) to make this clear:

"The formation of condensed organic products is commonly represented by lumped surrogate SVOCs in a 2-product model with volatilities and SVOC yields fitted to chamber experiments (Odum et al., 1996). To better represent the volatility of primary organic aerosol (POA) and the multigeneration oxidation of SVOCs to a wider range, Donahue et al. (2006) proposed the volatility basis set (VBS) model in which the mass yields of SVOCs are fitted to a fixed number of volatility bins (usually 0.01-105  $\mu$ g m-3). The VBS model has been adopted by several CTMs (such as WRF-Chem, GEOS-Chem, etc.)."

**Comment 2:** Page 6, line 162: Please add the appropriate references to support the nonvolatile nature of the products by these oxidation pathways.

**Response 2:** Formation of these non-volatile SOA was traditionally treated in the CMAQ model, except for dicarbonyls, IEPOX, and MAE that are assumed to form SOA by irreversible reactive uptake in our model, as an upper-limit estimation of SOA from these precursors. The text has been revised to make this clear (L162-173 in the manuscript):

"The SOA module mostly follows Pankow et al. (2015). Two types of SOA as traditionally treated in CMAQ were considered, "semi-volatile" (SV) portion that formed via equilibrium absorption-partitioning of SVOCs, and "non-volatile" (NV) portion that includes the oligomers and SOA formed via direct oxidation of aromatics at low-NOx. SOA from dicarbonyls, IEPOX, and MAE were formed by irreversible reactive uptake and categorized as NV-SOA in the current model as well. Some studies investigated SOA from glyoxal, methylglyoxal, and IEPOX using detailed reactions and reversible pathways in models or observed as reversible processes in chamber experiments, leading to a relatively lower SOA yield compared to the surface-controlled irreversible uptake (Lim et al., 2013; Knote et al., 2014; Galloway et al., 2009; EI-Sayed et al., 2018; Budisulistiorini et al., 2017). The non-volatile assumption used in this paper allows an upper-limit estimation of the importance of these additional SOA formation pathways."

**Comment 3:** Page 6 line 149: There is no discussion in the methodology about the observations and the statistical analysis metrics used to evaluate the model performance. Especially for the OA observations, there is no reference provided or description of the methods used.

**Response 3:** Thank you for pointing this out. We have included details of observation data, statistical analysis metrics for both meteorology and aerosols in the revised manuscript:

In L265-276

"The meteorological inputs and emissions have been used in several previous publications. Model performance on meteorological parameters (temperature and RH), gaseous species and gas and aerosol concentrations have been extensively evaluated (Hu et al., 2016; Hu et al., 2017; Qiao et al., 2018; Shi et al., 2017). A summary of the model performance related to this study is provided below. Observed meteorological data were obtained from the National Climatic Data Center (ftp://ftp.ncdc.noaa.gov/pub/data/noaa). Observations of OC at two urban locations, Beijing (Cao et al., 2014; Wang et al., 2015) and Guangzhou (Lai et al., 2016) and OA in Beijing (Sun et al., 2014) during January of 2013 as well as surface PM<sub>2.5</sub> at several monitoring sites during July of 2013 from China National Environmental Monitoring Center (http://113.108.142.147:20035/emcpublish/) were used to evaluate model estimates of aerosols. Details of measurement methodology and uncertainties of observations are listed in the corresponding references."

#### In L278-282

"Table 1 lists model statistics of mean observation (OBS), mean prediction (PRE), mean bias (MB), gross error (GE) and correlation coefficient (R) based on WRF and observations at monitoring sites located in 8 subregions of the domain (Figure S1) during January and July of 2013. The benchmarks for the MM5 model (another meteorology model) of 4-12km horizontal resolution suggested by Emery et al. (2001) are also listed in the table."

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In L294-296
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"Overall, the mean fractional bias (MFB) and mean fractional error (MFE) of OC are -0.20 and 0.27, within the criteria (MFB $\leq$ ±0.6; MFE $\leq$ 0.75) suggested by EPA (2007)."

#### In L308-309:

"The overall MFB and MFE of OA during January are -0.28 and 0.54, within the criteria (MFB≤±0.6; MFE≤0.75) suggested by EPA (2007)."

**Comment 4:** Page 7 line 179: More details are needed here. (1) How the model defines the low NOx and high NOx conditions? Which compounds each of the lump species represent? What is the difference between the lumped species of the same precursor (e.g., BNZ1, BNZ2, BNZ3)? (2) Can you include the aerosol yields for each lumped species in tables S1 and S2? Are these aerosol yields NOx-dependent? (3) What does the SVP stands for in Tables S1 and S2?

**Response 4:** (1) SOA formation in CMAQv5.0.2 is based on the frame of a previous version 4.7.1. All the details about "high" and "low"  $NO_x$  conditions (based on chamber experiments of corresponding VOCs), lumping species and method of each precursor, and the yields of precursors from parent VOCs have been documented by Carlton et al. (2010) and summarized in the revised supplemental materials as following:

"The CMAQ model treats high and low NOx SOA formation pathways during OH oxidation by allowing the lumped RO2 radical to competitively react with HO2 and NO. Using the lumped ARO1 species as an example, an SOA formation specific RO2 radical ARO1RO2 is added as a gas phase reaction product with OH:

ARO1 + OH  $\rightarrow$  ARO1RO2 + products

The ARO1RO2 can react with both HO2 and NO, as shown in the following two reactions:

 $ARO1RO2 + HO2 \rightarrow HO2 + TOLNRXN; k1$ 

 $ARO1RO2 + NO \rightarrow NO + TOLHRXN; k2$ 

Details of the determination of the rate constants can be found in Carlton et a. (2010). The TOLNRXN and TOLHRXN are counter species that track how much ARO1 is reacted through low NOx and high NOx pathways, respectively, in one gas chemistry time step. The concentrations of these counter species are passed into the aerosol module to calculate the formation semi-volatile products (TOL1 and TOL2) in the high NOx pathway and non-volatile products (TOL3) in the low NOx pathway, using the mass-specific yields, as listed in Table S1 and S2. Equilibrium partitioning of TOL1 and TOL2 in the gas phase and their counterparts ATOL1 and ATOL2 in the organic phase are affected by temperature and the amount of absorbing organics in the aerosol phase. Similar treatments are applied to the other lumped aromatic compounds ARO2, with xylene as a representative and most abundant species in that group, and to benzene. SOA formation from lumped long-chain alkene species ALK5, and isoprene and monoterpenes are not considered as  $NO_x$  dependent and are represented by equilibrium partitioning of one or two semi-volatile oxidation products. Details of the mass-specific yields of semi-volatile products and other related parameters can be found in Table S1 and S2."

We revised the text in (L188-190) to make it clear:

"More details about the lumped precursors such as formation conditions ("high" and "low" NOx), lumping species and method, and yields from parent VOCs can be found in Carlton et al. (2010) and summarized in SI."

(2) In the CMAQ model, the amount of SOA can form after a precursor reacts with OH,  $O_3$  or  $NO_3$  depends on the volatility of the products, which is temperature dependent, and the amount of the absorbing organics. The mass yields of the semi-volatile or non-volatile products are included in Table S1 and S2 in the revised supplementary materials. For more details, we refer the readers to Carlton et al. (2010) and the references therein.

(3) SVP is the saturation vapor pressure. We have explained this in the corresponding tables.

**Comment 5:** Page 8, lines 194-198: (1) Does the absorbing phase of equation 1 includes only the water associated with the organics (from eq. 3) or it includes the total water (including the water associated with the inorganic aerosol components)? (2) Under high RH (higher than the organic/inorganic phase separation RH, SRH), the aerosol organic phase is well mixed with the inorganic salts and, therefore, the aerosol water associated with the inorganic constituents can also contribute to the SOA absorbing medium (Pye et al., 2017). Please clarify what you have assumed here and add the relative discussion.

**Response 5:** (1) The absorbing phase of equation 1 only includes organic and water associated with organics when considering water-organic interactions. We explained this in the original text L195-196:

"In addition to organic compounds, water partitioning into OPM is enabled according to Eq 1 and Eq 2. In such a case, the absorbing phase in Eq 1 includes both organic aerosols and water partitioning into OPM."

(2) We didn't consider the mixing of organic and inorganic phase in this study and assumed that they are always two distinct aerosol phases without direct interactions. The phase separation RH (SRH) depends on the OM/OC ratio of the organic phase. Unlike the conditions modeled by Pye et al. (2017) for the southeast US where SOA is often the dominant OA component, the winter episode we modeled is dominated by primary emitted organic aerosols thus with a relatively low OM/OC (~1.4-1.6). The SRH based on equation (7) of Pye et al. (2017) is ~97%-99%. The summer episode has more contributions of SOA to OA, with OM/OC~1.8, which corresponds to an SRH of 87%. Thus, we don't expect interactions of organic and inorganic phases to occur in high frequency to greatly influence the model results. We assumed no interactions between inorganic and organic phases in the current model. We have also revised text in L207-208:

"In the current model, we assumed no interactions between the inorganic and organic phases."

**Comment 6:** Page 8, Equation 3: This equation gives the volume of water associated with the organic fraction of the aerosol. However, ALW on the left-hand side refers to the mass of water. Please correct.

**Response 6:** The equation has been corrected in the revised text L212-214: "

$$ALW_{org} = \rho_w V_{org} \kappa_{org} \frac{a_w}{1 - a_w}$$
(Eq3)

where  $\rho_w$  is the density of water (assumed to be 1 g cm<sup>-1</sup>), V<sub>org</sub> is the volume concentration of organics, and  $a_w$  is the water activity (assumed to be the same as RH)."

**Comment 7:** Page 8, lines 203-204: How do you estimate the hygroscopicity? **Response 7:** We used Eq3 to estimate the hygroscopicity. This has been clarified in the revised text in L214-216:

"Since ALW<sub>org</sub> in this study is calculated mechanistically using the partitioning theory,  $\kappa_{org}$  can be estimated by rearranging Eq3:

$$\kappa_{org} = \frac{ALW_{org}}{\rho_w V_{org}} \times \frac{1 - a_w}{a_w}$$
(Eq4)

**Comment 8:** Page 8, lines 204-205: How do you calculate the ALW? Are you using  $\kappa_{org}$  and the eq3? Do you use the kappa hygroscopicity to calculate the

ALW? If so, how you estimate the kappa?

"

**Response 8:** In this study, the ALW<sub>org</sub> is independently calculated by mechanistically allowing water molecules to partition into the organic phase with UNIFAC calculated activity. In such a case, Eq 3 can be used to provide an independent estimation of  $\kappa_{org}$ . Linear regression analysis can be performed using the calculated  $\kappa_{org}$  against the model calculated O/C ratios, as shown in Figure 4. We have removed this sentence to avoid confusion.

**Comment 9:** Page 9, lines 212-216: Add a reference to tables S1 and S2. Furthermore, why the values of OM:OC in tables S1/S2 are different than the values provided by your reference (Pye et al., 2017)?

**Response 9:** We used OM:OC ratio in Pankow et al. (2015). There were mistakes in the original Table S1 and S2 and have been corrected in the revised supplemental materials. The text has also been revised in L228-229 as following:

"The OM:OC ratio of each SOA component follows Pankow et al. (2015) as shown in Table S1-S2."

**Comment 10:** Page 9 lines 217-218: Can you add the total size of the model domain?

**Response 10:** The text has been revised in L231-233 as follows:

"The simulation domain has a horizontal resolution 36 km × 36 km (100  $\times$  100 grids) and a vertical structure of 18 layers up to 21 km, which covers eastern China as shown in Figure S1."

**Comment 11:** Page 9 lines 233: (1) What boundary conditions are used? Please make a comment on how these can affect the simulation results. (2) I would recommend adding spatial maps of primary organic aerosol emissions and SOA precursor emissions and summarizing in a Table the domain average emission rates of POA and each SOA precursors.

**Response 11**: (1) We used a predefined boundary profile in CMAQ that represents a clean continental condition. The northern and western boundaries, as well as areas to the further north and west, are mostly remote areas with much lower emissions. The mountains in the north and west part of the domain also limit the influence of emissions enter from the boundaries to the central part of the domain. The influence of marine air from the south and east boundaries is also small, as local emissions dominate the concentrations.

(2) A figure and a table showing emissions of POA and SOA precursors were added to the revised supplemental materials as Figure S2 and Table S4.

**Comment 12:** Pages 9-10 lines 234-238: (1) This paragraph needs to be expanded and written in a separate section. In this section the authors should describe in more detail the following: i) Basecase simulation. Please explain how the default CMAQ is simulating POA and SOA and how different is this modelling configuration with the one the authors are testing, ii) Sensitivity simulations. Please explain in much more detail the sensitivity simulations conducted in this work. (2) In addition the authors say that they have conducted three sensitivity scenarios named S1, S2 and S3. In their manuscript they only show results from S3 and they never discuss the results of S1 and S2.

**Response 12:** (1) Details of how default CMAQ simulates SOA formation through equilibrium partitioning of lumped semi-volatile products into the

organic phase (which include both POA and SOA) have been described by Carlton et al. (2010) and Hu et al. (2017) so we don't think it is necessary to repeat it here. We expanded each simulation scenario in details in the revised text in L249-259:

"Four scenarios are investigated in this study. The base case (BS) applies the default secondary organic aerosol module of CMAQ v5.0.1. In this case, no water partitioning into OPM is considered. Lumped semi-volatile products from the oxidation of various precursors partition into a single organic phase, which is considered as an ideal mixture of POA and SOA with  $\gamma_{org}$ =1. The water case (S1) includes water partitioning into OPM, which is again considered as an ideal solution ( $\gamma_{org}$ =1 and  $\gamma_{H2O}$ =1). The UNIFAC case (S2) considers the interaction between organic constituents with UNIFAC calculated activity coefficients ( $\gamma_{org}$ ≠1) but does not allow water partitioning into OPM. The combined case (S3) allows both water partitioning and interactions between all constituents (including water and organics) using UNIFAC calculated activity coefficients ( $\gamma_{org}$ ≠1 and  $\gamma_{H2O}$ ≠1). "

(2) The impacts of S1 and S2 were discussed in Section 4 of the original manuscript Page 16 L405-419. We also included a description of the outlines of the results and discussions in the original manuscript Page 6 L140-148. The text has been revised in L259-262 to make it clear:

"The results of BS and S3 are used to examine the overall impacts of water partitioning into OPM and polarity of organics on SOA and ALWorg, as shown in Section 3.1-3.4. The separate influences of those two processes on SOA from S1 and S2 are discussed in Section 3.5."

**Comment 13:** Page 10 line 240: The section "Model evaluation" is extremely problematic and raises guestions on the validity of the modelling results given that the model evaluation is insufficient. More specifically: 1) Given that the CMAQ default configuration has been modified to consider the importance of water and organic compound polarity on SOA formation, an accurate evaluation of the updated model performance is needed. 2) You should compare the model results for organic compounds during both July and January 2013. Currently, the evaluation includes a comparison with OC observations only during January over only three locations of the relatively large model domain. Furthermore, please mention in the text what factor have you used to convert the modeled OA to OC. 3) The total PM2.5 measurements have been used to evaluate the model performance during July without explaining the rationale of this choice since the focus of this study is solely the organic fraction of the aerosols. I suggest removing the PM2.5 evaluation or at least moving it to the supplement. 4) Can you include more OC/OA observations over other areas of their domain in your evaluation? 5) It is also important to compare the simulated POA and SOA against observations (e.g., from AMS). Furthermore, it would be helpful to show how the model performance against SOA measurement changes between the BC and the S1, S2, S3 cases.

**Response 13:** (1) For the updates in CMAQ except for water partition into OPM and non-ideality of the organic-water mixture, previous studies have extensively examined the model performances and will not be further discussed in detail in this work. The text has been revised in L149-161 to reflect this:

"Model configurations were largely based on that used by Hu et al. (2016) as summarized below. Firstly, SAPRC-11 was expanded for a more detailed treatment of isoprene oxidation and tracking dicarbonyl products (glyoxal and methylglyoxal) from different groups of major precursors (Ying et al., 2015). Secondly, SOA from isoprene epoxydiols (IEPOX), methacrylic acid epoxide (MAE) and dicarbonyls through surfacecontrolled irreversible reactive uptake were added (Hu et al., 2017; Li et al., 2015; Liu et al., 2020; Ying et al., 2015). Thirdly, the heterogeneous formation of secondary nitrate and sulfate from NO<sub>2</sub> and SO<sub>2</sub> reaction on the particle surfaces (Ying et al., 2014) were added, which is an important source of secondary inorganic aerosols (Zheng et al., 2015) and improves model estimates of nitrate and sulfate (Qiao et al., 2018; Shi et al., 2017). Fourthly, SOA yields were corrected for vapor wall loss (Zhang et al., 2014). Impacts of the above updates on model performances have been extensively discussed in the cited work and will not be further investigated in the current study."

(2) Unfortunately, detailed chemical composition measurements for aerosols are very limited in China during 2013. We only have observations of OC and OA in January of 2013 and PM<sub>2.5</sub> in July of 2013 available for model evaluation. Thus, we opt to use the most relevant data to provide a very limited assessment of the capability of the model in predicting SOA. The factors for OA to OC conversion follow the OM:OC ratio listed in Table S1 and S2 for SOA. POC is directly predicted by the model. The text has been revised in L291-292 to make this clear:

### "The factors used to convert SOA to OC (SOC) are listed in Table S1-S2. OC from POA (POC) is directly predicted by the model."

(3) Even though limited OC/OA measurements are available to us during this period, the base case model is later applied by another research group to model wintertime SOA formation in east China (Liu et al., 2020). The predicted OC and SOA agree well with observations (Figure 2 of Liu et al. 2020), and the model performance statistics for OC and SOA are similar to those of PM<sub>2.5</sub>. We agree with the reviewer that  $PM_{2.5}$  is not an ideal indicator to evaluate the capability of the model in predicting SOA, however, as a significant fraction of  $PM_{2.5}$  in July is secondary, this still provides an indirect assessment of the model prediction of oxidation capacity of the atmosphere, which is import for

SOA formation. Additional modeling studies are needed to evaluate the performance of the model in summer.

We have explained this in the revised text (L327-329) as following:

"Due to the lack of observed OC and OA in July of 2013, as an alternative, model performances are evaluated by comparing predicted and observed PM<sub>2.5</sub> at ground sites (Figure S1) as shown in Figure S3."

(4) We do not have more OC and OA data for the simulating episode of 2013.

(5) We did not have SOA observations in this episode. We compared the modeled SOA/POA ratio with AMS observations from other literature, finding a significant underestimation in the current model. This bias might be due to missing SOA from combustion (intermediate volatile organic compounds, IVOCs) and not treating POA as semi-volatile. We added a discussion about the model bias in the revised manuscript in L309-326:

"Again, no apparent changes of SOA nor OA are observed between case S3 and BS (not shown), since POA is predicted to be the primary contributor to OA at Beijing in winter in the current model, with an averaged SOA/POA ratio of 0.12. This ratio is much lower than the field observation of about 0.45-1.94 (Zhao et al., 2019; Sun et al., 2013; Sun et al., 2016). The bias might be due to the missing SOA converted by partitioning and aging of semi-volatile POA as well as oxidation from intermediate volatile organic compounds (IVOCs) and VOC oxidation products. Those pathways are shown to be important for SOA formation by modeling, field and chamber studies (Hodzic et al., 2010; Jimenez et al., 2009; Murphy et al., 2017; Robinson et al., 2007; Shrivastava et al., 2008; Tkacik et al., 2012; Zhao et al., 2014; Zhao et al., 2016a).

A sensitivity test was performed by using the newest CMAQ model version 5.3.1 that includes all the above processes in the aerosol module. The SOA/POA ratio in Beijing is improved greatly to be 0.83 in winter. However, high uncertainties still exist in the emissions of the involved precursors and characterization of SOA formation through these processes, needing further constrains by observations. Their influences on water partitioning into OPM and non-ideality of the organic-water mixture on SOA will be evaluated in a future study."

Since SOA is underestimated, no significant differences in BS and S3 are observed. Case S1 and S2 are designed for the sensitivity test of model results to water partitioning and non-ideality of condensed organics separately. Therefore, we did not evaluate model performances from S1 and S2.

**Comment 14:** Page 10 lines 251-253: The authors state here that the impacts of water-cocondensation and polarity of organic condensed species on SOA

formation are not significant during winter. This highlights the need to evaluate their model results during July where they have found significant changes with the basecase simulation. Furthermore, the results from the three sensitivity simulations should be evaluated individually.

**Response 14**: Unfortunately, we have no observations of OC, OA or (and) SOA of July 2013. We tried to evaluate model performances by comparing the predicted and observed  $PM_{2.5}$  as an alternative. This has been explained in the revised manuscript in L327-329:

"Due to the lack of observed OC and OA in July of 2013, as an alternative, model performances are evaluated by comparing predicted and observed PM2.5 at ground sites (Figure S1) as shown in Figure S3."

Since SOA is underestimated, no significant differences in BS and S3 are observed. Case S1 and S2 are designed for the sensitivity test of model results to water partitioning and non-ideality of condensed organics separately. Therefore, we did not evaluate model performances from S1 and S2.

**Comment 15:** Page 10 lines 254-256: These are indeed possible factors. Can the authors comment, based on their analysis, which of these two possible factors is more important and try to be more specific? A comparison with AMS observations would be helpful here.

**Response 15:** We did a sensitivity test by simulating the same episode with the current CMAQv5.3.1 in which POA was treated as semi-volatile and aging in the gas phase. Also, a missing source of SOA from intermediate VOCs (IVOCs) oxidation and aging of IVOCs and VOCs oxidation products (pcSOA) was added in CMAQv5.3.1. The results showed that the modeled SOA/POA has been improved from 0.12 of case S3 to 0.83 (Figure R1), more close to AMS observations. However, there were still some peak values underestimated by the model, which might be due to the uncertainties of POA emissions. We have revised the text (L309-326), as mentioned in Response 13 (5).



**Figure R1.** Modeled concentration of semi-volatile POA (sv-POA), OA and fraction of each organic aerosol component fsvPOA (sv-POA), foxPOA (oxidation of sv-POA), fpcSOA (pc-SOA) and fsvSOA (traditional SOA) in Beijing (a, c) and Guangzhou (b, d). Observations of OA in January 2013 at

Beijing (Obs.) are also included in (a). The left axis is the concentration ( $\mu$ g m<sup>-</sup>

<sup>3</sup>) and the right axis is the fraction of OA components.

**Comment 16:** Page 11 line 263: The aging of POA, under specific conditions can enhance the SOA formation, especially over polluted areas. Can the authors comment how this important omission of their model configuration can affect their result? Once again, a comparison against POA and SOA from AMS observations will be helpful to identify the limitations of their model due to the treatment of POA as non-volatile and non-reactive.

**Response 16:** We did a sensitivity test by CMAQv5.3.1 that includes SOA from POA aging, IVOCs oxidation, and aging of IVOCs and VOCs oxidation products. The modeled SOA/POA is improved greatly in Beijing with no significant improvement in OA compared to the results of case S3. A significant improvement of SOA was observed from the contribution from IVOCs oxidation, and the aging of IVOCs and VOCs oxidation products (pc-SOA). Discussions about this have been included in the revised text in L309-326, as mentioned in Response 15.

Also, we examined the sensitivity of SOA and organic liquid water ( $ALW_{org}$ ) to pcSOA and POA in an offline calculation in Beijing, Guangzhou, Jinan, and Nanjing. POA has the same properties as we used in the model. Non-volatile isoprene SOA is taken to represent pcSOA as their similarities in saturation vapor pressure and O:C ratio. We found that both SOA and  $ALW_{org}$  are

positively correlated with pcSOA, increased by 2-5 times in different locations when pcSOA increased by 2 times.

**Comment 17:** Page 11 line 264: Please provide two spatial maps of the fraction SOA/Total OA during January and July 2013 so as to show the contribution of POA and SOA to total OA during each simulation period.

**Response 17:** We have added a figure of SOA/OA ratio by case BS and S3 (as shown below) in the revised supplemental materials (Figure S13).



**Figure R2**. Averaged SOA/OA ratio from case BS and S3 during January and July of 2013.

**Comment 18:** Page 13 line 317: You need a zonal map to show how the water partition changes with altitude and not the total column.

**Response 18:** Most of the SOA and ALW<sub>org</sub> retain in the lower levels of the troposphere. The information on altitude variation may not be very useful. Thus, no changes were made regarding this comment. We have also deleted this sentence in the revised manuscript to avoid confusion.

**Comment 19:** Page 13, lines 322-336: It is not clear how you calculate the  $\kappa$ 

*org in your model. This is very important for this section.* **Response 19:** We explained this in the revised text in L214-216: "Since ALWorg in this study is calculated mechanistically using the

partitioning theory,  $\kappa_{org}$  can be estimated by rearranging Eq3:

"

$$\kappa_{org} = \frac{ALW_{org}}{\rho_w V_{org}} \times \frac{1 - a_w}{a_w}$$
(Eq4)

**Comment 20:** Figure 1: The quality of the figure is poor. It is extremely difficult to see all the plotting data and the changes due to the use of different scenarios (especially in figure 1b).

**Response 20:** We removed the results of BS since they are very similar to those of S3. Now the figure has been revised as follows:



**Figure 1.** Comparison of (a) observed and modeled organic carbon concentration at University of Beihang (BH), Tsinghua University (TH) and Guangzhou (GZ); (b) observed organic aerosol (Obs-OA) at Beijing and predictions of total OA (pOA) and SOA (pSOA), unit is  $\mu$ g m-3. Locations of monitoring sites are shown in Figure S1.

Comment 21: Figure 2: I found the use of daily maximum concentration in the

"difference" maps misleading. Given that you have the monthly average SOA from the basecase simulation, I would prefer to see the absolute (and relative) change of the monthly average SOA due to the use of S3 as well, and not the daily maximum. Furthermore, please add (a), (b), etc. to each subplot of the figure and add this information to the figure caption (apply this change in the rest of the figures as well).

**Response 21:** The monthly-averaged daily maximum differences have been replaced by the monthly-averaged differences for all the corresponding figures to reflect the general impacts on SOA and ALW<sub>org</sub>. Each panel of the figure is labeled in sequence.

**Comment 22:** Figure 4: All the fitted correlations listed here suggest that compounds with very low or zero O:C have negative hygroscopicity. Can you comment on this limitation and include a discussion in the text?

**Response 22:** The relatively low values of hygroscopicity for low O:C ratio might be due to the linear regression. We also did an exponential fitting for the two variables so that the hygroscopicity falls in the range of (0,1) and is positively correlated with O:C ratio. The text and figure 4 have been revised accordingly.

In L402-407:

"In both seasons,  $\kappa_{org}$  approaches zero and negative values as O:C decreases, which might be due to the linear regression of  $\kappa_{org}$  and O:C. To avoid this, an exponential fitting of the two variables is performed so that  $\kappa_{org}$  falls in the range of (0,1) and is positively correlated with O:C. In this case, the fitted correlations are  $\kappa_{org}$ =1-exp(-(O:C/1.88)<sup>2.29</sup>) and  $\kappa_{org}$ =1-exp(-(O:C/1.06)<sup>4.50</sup>) for January and July of 2013, respectively."



**Figure 4**. The correlation of hygroscopicity of organic aerosol ( $\kappa_{org}$ ) and O:C ratio at 9 representative cities including Shenyang (SS), Beijing (BJ), Jinan (JN), Zhengzhou (ZZ), Xi'an (XA), Nanjing (NJ), Shanghai (SH), Chengdu (CD), and

Guangzhou (GZ) in January (a) and July (b) of 2013. O:C ratios are categorized into 10 bins. In each bin, the ranges of O:C and  $\kappa_{org}$  are represented by bars. The mean values of O:C and  $\kappa_{org}$  are represented by triangles colored by the averaged RH of each bin. The relationship between  $\kappa_{org}$  and O:C is fitted by a linear function with reduced major axis regression (blue lines) and an exponential function (red lines), respectively.  $\kappa_{01}$  and  $\kappa_{07}$  represent the fitted correlation for January and July, respectively.

**Comment 23:** Table S3. Please explain in the table what the fraction in the last column stands for. How have you estimated the molecular weight and fraction of the POA from unknown compounds?

**Response 23:** The last column is the molar fraction of each POA surrogate. We have clarified this in the first row of this column in the table. The molecular weight and molar fraction of unknown compounds of POA have been listed in the original table already, which are 390 and 0.3, respectively.

### **Minor Comments**

**Comment 24:** The language and structure of the sentences can be substantially improved in many parts of the manuscript. Just a few examples are listed here, but I suggest revising thoroughly the wording in the whole text. **Response 24:** We thank the reviewer for pointing this out. The whole text, as well as figures and tables, have been revised carefully.

**Comment 25:** Page 5, lines109-11: Please rephrase. The sentence sounds wrong.

**Response 25:** The text has been revised in L109-111 as follows:

"Laboratory and field studies have observed water absorbed by SOA from a variety of precursor VOCs (Lambe et al., 2011; Zhao et al., 2016b; Asa-Awuku et al., 2010; Varutbangkul et al., 2006)."

Comment 26: Page 5, lines117-118: Please rephrase.

**Response 26:** The sentence has been revised in L116-117 as follows: "The total water content is the summation of water associated with each solute at the same water activity."

**Comment 27:** Page 6 line 144: OC and OA abbreviations have not been used before in the main text.

**Response 27:** OA abbreviation has been explained in a previous part of the revised manuscript in L120-122:

"Pye et al. (2017) found that the modeled organic aerosol (OA) improved significantly but biased high at nighttime when ALW<sub>org</sub> is included in the calculation."

Thus, this sentence has been revised in L139-141 as follows:

"The model performance was evaluated against observed meteorological parameters (temperature and relative humidity, RH) as well as PM<sub>2.5</sub>, organic carbon (OC), and OA at ground monitoring sites."

**Comment 28:** Page 8, line 204: Please correct the "can estimated" to "can be estimated".

**Response 28:** The text has been revised as instructed.

**Comment 29:** Page 8 line 206: I would use the word "correlate" instead of "dependent"

**Response 29:** The text has been revised in L219-229 as follows:

"In many studies,  $\kappa_{ora}$  is assumed to increase linearly with the oxidation

state of OA, expressed as the O:C ratio (Massoli et al., 2010;Duplissy et al., 2011;Lambe et al., 2011)."

**Comment 30:** *Page 9 line 229: Change "in" with "on"* **Response 30:** The text has been revised as instructed.

**Comment 31:** Page 10 line 254: Change the sentence to: "In Beijing and Guangzhou, these impacts are not significant during winter"

**Response 31:** The sentence has been revised in L298-300 as:

"No significant differences in OC are observed in S3 compared to BS (not shown), likely due to the biased-low SOA predicted in the current model so that limits the impact of ALW<sub>org</sub> on SOA formation."

Comment 32: Page 11 line 274-275: Please rephrase. Response 32: The sentence has been revised as follows: "The criteria of MFB and MFE followed recommendations by Boylan and Russell (2006)."

**Comment 33:** Page 11 line 282: Which two areas? You have mentioned several areas in the previous sentence.

**Response 33:** The sentence has been revised as follows:

"Monthly-averaged SOA concentrations in the above areas are up to 25 and 15-20  $\mu$ g m<sup>-3</sup>, respectively."

Comment 34: Page 13 line 317: "column water". Please rephrase Response 34: The text has been revised in L378 as follows: "Based on the column concentrations of ALW<sub>org</sub> and ALW<sub>org</sub>/SOA ratio (Figure S8),"

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## 1 Impacts of water partitioning and polarity of organic compounds on

## 2 secondary organic aerosol over Eastern China

3 Jingyi Li<sup>1, 2</sup>, Haowen Zhang<sup>2</sup>, Qi Ying<sup>3,\*</sup>, Zhijun Wu<sup>4, 1</sup>, Yanli Zhang<sup>5,6</sup>, Xinming

- 4 Wang<sup>5,6,7</sup>, Xinghua Li<sup>8</sup>, Yele Sun<sup>9</sup>, Min Hu<sup>4, 1</sup>, Yuanhang Zhang<sup>4, 1</sup>, Jianlin Hu<sup>1, 2, \*</sup>
- <sup>5</sup><sub>6</sub> <sup>1</sup> Collaborative Innovation Center of Atmospheric Environment and Equipment

Technology, Nanjing University of Information Science & Technology, Nanjing 210044,

- 6 China
- <sup>9</sup> <sup>2</sup> Jiangsu Key Laboratory of Atmospheric Environment Monitoring and Pollution
- Control, School of Environmental Science and Engineering, Nanjing University of Information Science & Technology, Nanjing 210044, China
- <sup>3</sup> Texas A&M University, College Station, Texas 77843, USA
- <sup>4</sup> State Key Joint Laboratory of Environmental Simulation and Pollution Control, College
- 13 of Environmental Sciences and Engineering, Peking University, Beijing 100871, China
- <sup>5</sup> State Key Laboratory of Organic Geochemistry and Guangdong Key Laboratory of
- 15 Environmental Protection and Resources Utilization, Guangzhou Institute of
- 16 Geochemistry, Chinese Academy of Sciences, Guangzhou 510640, China
- <sup>6</sup>Center for Excellence in Regional Atmospheric Environment, Institute of Urban
- 18 Environment, Chinese Academy of Sciences, Xiamen 361021, China
- <sup>19</sup> <sup>7</sup> University of Chinese Academy of Sciences, Beijing 100049, China
- <sup>8</sup> School of Space & Environment, Beihang University, Beijing 100191, China
- <sup>9</sup> State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric
- 22 Chemistry, Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing
- 23 100029, China
- 24
- 25 Corresponding authors:
- 26 Qi Ying, Email: qying@civil.tamu.edu
- 27 Jianlin Hu, Email: jianlinhu@nuist.edu.cn
- 28

#### 29 Abstract

Secondary organic aerosol (SOA) is an important component of fine particular matter 30 31 (PM<sub>2.5</sub>) in China). Most air quality models use an equilibrium partitioning method along 32 with estimated the saturation vapor pressure (SVP) of semi-volatile organic compounds (SVOCs) to predict SOA formation. However, this method ignores the models typically 33 34 assume that the organic particulate matter (OPM) is an ideal mixture and ignore the 35 partitioning of water vapor to the organic aerosols and the organic phase non ideality, both of which affect the partitioning of SVOCs.OPM. In this study, the Community Multi-scale 36 Air Quality model (CMAQv5.0.1) was used CMAQ) is updated to investigate the above 37 impacts impacts of water vapor partitioning and non-ideality of the organic-water mixture 38 39 on SOA formation during winter (January) and summer (July) of 2013 over eastern China. 40 The organic aerosol module was updated by incorporating watermodel treats the 41 partitioning of water vapor molecules into OPM and uses the UNIFAC model to estimate the activity coefficients of species in the organic particulate matter (OPM) and considering 42 non-ideality of organic-water mixture. The modified model can generally capture the 43 observed surface organic carbon (OC), the total organic aerosol () with a correlation 44 coefficient R of 0.7, and the surface OA) and diurnal variation of PM<sub>2.5</sub> at ground sites. 45 with the mean fractional bias (MFB) and mean fractional error (MFE) of -0.28 and 0.54, 46 respectively. SOA concentration shows significant seasonal and spatial variations, with 47 high concentration levels concentrations in the North China Plain (NCP), Central China and 48 Sichuan basin (SCB) areas regions during winter (up to 25 µg m<sup>-3</sup>) and in the Yangtze River 49 Delta (YRD) during summer (up to 1216 µg m<sup>-3</sup>). When water partitioning is included in 50 In winter, SOA concentrations increase decreases slightly in the updated model, with the 51 monthly-averaged daily maximum relative difference change of 10-20% in the highly 52 concentrated areas, mainly due to organic-water interactions. The monthly-averaged 53 54 concentration of SOA increases greatly in summer, by 20-50% at the surface and 10-30% for-60% in the whole column, mostly. The increase of SOA is mainly due to the increase 55
in biogenic SOA in inland areas and anthropogenic SOA. The increase in SOA is more 56 significant in summer, by 20-90% at the surface and 30-70% for the whole column. The 57 increase of SOA over the land is mostly due to biogenic SOA while the increase of SOA 58 59 over the in coastal regions is related with that of anthropogenic origin. Further analysis of two representative cities, Jinan and Nanjing, shows that changes of SOA are favored 60 under hot and humid conditions. The increases in SOA cause a 12% elevation in theareas. 61 62 As a result, the averaged aerosol optical depth (AOD) is increased by up to 10% and  $\frac{15\%}{100}$ 63 enhancement in the cooling effects effect of aerosol radiative forcing (ARF) is enhanced by up to 15% over YRD in summer. The aerosol liquid water content associated with OPM 64 65 (ALW<sub>org</sub>) at the surface is relatively high over the landin inland areas in winter and over 66 the ocean in summer, with the monthly-averaged daily maximum concentration of 2-90.5-3.0 and 5-127 µg m<sup>-3</sup>, respectively. By using the The hygroscopicity parameter ( $\kappa$ ) of OA 67 <u>based on the  $\kappa$ -Köhler theory, we calculated the hygroscopicity of OA with is</u> 68 determined using the modeled ALW<sub>org</sub>, finding that the. The correlation of  $\kappa$  with O:/C 69 70 ratio varies significantly across different cities and seasons. Water-Analysis of two representative cities, Jinan (in NCP) and Nanjing (in YRD), shows that the impacts of water 71 partitioning and non-ideality of the organic-water mixture on SOA are sensitive to 72 temperature, relative humidity (RH), and the SVP of SVOCs. The two processes exhibit 73 opposite impacts on SOA in eastern China. Water uptake increases SOA by up to 80% in 74 the organic phase, while including non-unity activity coefficients decreases SOA by up to 75 50%. Our results indicate that both water partitioning into OPM only promotes SOA 76 formation, while non-ideality of organic-water mixture only leads to decreases in SOA in 77 most regions of eastern China. Water partitioning into OPM and the activity coefficients of 78 the condensed organics should be considered in air quality models in simulating SOA 79 formation from gas-particle partitioning, especially in hot and humid environments. 80 81

82 Keywords: SOA, non-ideality, water partitioning, hygroscopicity

83

#### 84 **1 Introduction**

85 Secondary organic aerosol (SOA) is formed via a complex interaction of volatile organic compounds (VOCs) with oxidants and primary particles emitted from anthropogenic and 86 biogenic sources in the atmosphere. As an important component of fine particular matter 87 (PM<sub>2.5</sub>), SOA can cause severe air pollution in urban and suburban areas (Huang et al., 88 89 2014) and exhibit adverse health effects (Polichetti et al., 2009; Feng et al., 2016; Xing et 90 al., 2016; Atkinson et al., 2014). (Atkinson et al., 2014). SOA also plays an important role in new particle formation and particle growth (Man et al., 2015; Zhang et al., 91 2011; Wiedensohler et al., 2009; Yue et al., 2011; Liu et al., 2014; Ehn et al., 2014; Huang et 92 93 al., 2019; Jokinen et al., 2015) and further contributes that further contribute to the enhancement of cloud condensation nuclei (CCN) (Yue et al., 2011; Wiedensohler et al., 94 95 2009;Liu et al., 2014;Jokinen et al., 2015)(Wiedensohler et al., 2009; Ehn et al., 2014). 96 This will, in turn, impact the atmospheric aerosol burden, precipitation and water 97 circulation, solar radiation budget, and climate (Rosenfeld et al., 2008;Spracklen et al., 2011; Quaas et al., 2008; Ramanathan et al., 2001; Hatzianastassiou et al., 2007; Hegerl et al., 98 99 2015). However, the mechanisms of these (Ramanathan et al., 2001). However, the extents of those influences are not well understood so far, due to the high uncertainties associated 100 with the formation and physical and chemical properties of SOA (Shrivastava et al., 101 2017)(Shrivastava et al., 2017). Large gaps still exist in SOA mass loadingsloading and 102 103 properties between model estimates models and laboratory and field measurementsobservations (Gentner et al., 2017; Ervens et al., 2011; Hayes et al., 104 2015)(Gentner et al., 2017; Ervens et al., 2011; Hayes et al., 2015). Therefore, it is crucial 105 to explore and resolve this issue to improve our knowledge of the roles of SOA in the 106 107 environment, human health, and climate.

108Gas-particle partitioning of semi-volatile and low-volatile organic compounds109(SVOCs and LVOCs) generated from VOC oxidation is an important pathway of SOA

110 formation. In most current chemical transport models (CTMs), this process is treated as an 111 equilibrium partitioning process that depends on the mass concentration of the organic particulate matter (OPM), ambient temperature, (T), the mean molecular weight of OPM, 112 113 and the volatility of purer condensed organics (Pankow, 1994)(Pankow, 1994). The volatilities formation of condensed organic products from a certain precursor VOC are 114 either is commonly represented by that of several-lumped surrogates based on surrogate 115 116 SVOCs in a 2-product model with volatilities and SVOC yields fitted to chamber experiments (2-product model) (Odum et al., 1996)(Odum et al., 1996) or fitted into 117 different bins of a fixed volatility range (usually 0.01-10<sup>5</sup> µg m<sup>-3</sup>) (volatility basis set model, 118 VBS model) (Donahue et al., 2006). Although the above. To better represent the volatility 119 of primary organic aerosol (POA) and the multi-generation oxidation of SVOCs to a wider 120 range, Donahue et al. (2006) proposed the volatility basis set (VBS) model in which the 121 mass yields of SVOCs are fitted to a fixed number of volatility bins (usually  $0.01-10^5 \,\mu g$ 122  $m^{-3}$ ). The VBS model has been adopted by several CTMs (such as WRF-Chem, GEOS-123 124 Chem, etc.).

Although the models can capture the general trend of SOA evolution and mass 125 concentration to some extent (Slowik et al., 2010; Li et al., 2017a; Baek et al., 126 2011;Bergström et al., 2012;Woody et al., 2016;Heald et al., 2006)(Li et al., 2017a; 127 Bergström et al., 2012; Woody et al., 2016), both of them neglected, two key factors 128 that currently neglected in models may lead to biases: 1) the molecular structures and 129 interactions of functional groups (-OH, -C=O, -COOH, etc.) of condensed organics (non-130 131 ideality); 2) partitioning of water vapor, anthe most abundant atmospheric constituent besides  $O_2$  and  $N_2$ , to OPM. The non-ideality alters the volatility of condensed organics, 132 and thus their contributions to the total SOA mass loading (Cappa et al., 2008). Water 133 partitioning into OPM can reduce the partial pressure of organics and leaddue to increase 134 135 in SOA mass, which is called the Raoult's Law effect (Prisle et al., 2010)(Prisle et al., 2010). This impact and lead to increases in SOA mass. The amount of aerosol liquid water 136

associated with organics (ALW<sub>org</sub>) may vary for different SOA-precursors (Healy et al.,
2009;Prisle et al., 2010)(Healy et al., 2009; Prisle et al., 2010). The above two aspects will
not only affect the chemical composition of SOA but also the inorganic portion (Ansari
and Pandis, 2000;Meyer et al., 2009) and optical properties (Liu and Wang, 2010;Denjean
et al., 2015)(Ansari and Pandis, 2000) and optical properties (Denjean et al., 2015) of
aerosols.

143 Laboratory and field studies have confirmed the fact that observed water absorbed by 144 SOA (quantified as hygroscopicity,  $\kappa$ )-from a variety of precursor VOCs (Lambe et al., 2011;Zhao et al., 2016b; (Lambe et al., 2011; Zhao et al., 2016b; Asa-Awuku et al., 145 2010; Varutbangkul et al., 2006); Varutbangkul et al., 2006). The hygroscopicity of SOA-is 146 highly, quantitatively described by the hygroscopicity parameter,  $\kappa$ , is correlated with the 147 oxygen-to-carbon ratio (O:C) and increases with more oxidized SOA during photochemical 148 149 aging (Poulain et al., 2010; Wang et al., 2014; Lambe et al., 2011; Tritscher et al., 150 2011a;Zhao et al., 2016b;Massoli et al., 2010;Tritscher et al., 2011b;Duplissy et al., 2011)(Lambe et al., 2011;Zhao et al., 2016b). The OPM-associated water partitioning can 151 be estimated using the κ-Köhler theory under the Zdanovskii-Stokes-Robinson (ZSR) 152 153 assumption of no interactions between any constitutes in aerosols (Petters and Kreidenweis, <del>2007)</del>(Petters and Kreidenweis, 2007). The 154 total water content is the summarization summation of water associated with each constitute solute at the same 155 156 **RH.**water activity. Guo et al. (2015)(2015) found that this simplified method, along with 157 the ISORROPIA model which is used to predict aerosol liquid water (ALW) associated 158 with the inorganic portion of aerosols, reproduced could reproduce the observed total ALW 159 in the ambient environment. Pye et al.  $\frac{(2017)}{(2017)}$  applied this approach along with a parameterization of overall  $\kappa$  based on O:C ratio and a simplified method to estimate 160 activity coefficients of organics and found that the modeled organic aerosol (OA and 161 162 ALW are) improved during daytimesignificantly but still-biased lowhigh at nighttime-Shortcomings still exist in the above method for water associated organics ( when ALW<sub>org</sub>) 163

164 is included in the calculation. However, as interactions between the interaction among organic species and between organics and water in the organic-water mixture are not 165 166 considered, which has been shown to play an important role in SOA formation and water 167 partitioning to OPM (Kim et al., 2019)(Kim et al., 2019). A representation of water partitioning along with SVOCs with consideration of water-organic and organic-organic 168 interactions in CTMs showed significant influences in SOA and ALW in the eastern U.S. 169 170 where biogenic SOA dominated in OA and the internal mixing assumed for the aerosol 171 (Pankow et al., 2015; Jathar et al., 2016).

China has been suffering from severe PM2.5 pollution especially in the eastern region 172 173 with fast urbanization and economic development, the ALW<sub>org</sub> estimated by the  $\kappa$ -Köhler theory and its impact on SOA might not be accurate. Using UNIversal Functional Activity 174 175 Coefficient (UNIFAC) method (Fredenslund et al., 1975) for calculating activity 176 coefficients of the organic-water mixture, it was found that in the eastern U.S., where biogenic SOA dominated the OA, considering ALW<sub>org</sub> leads to a significant increase in 177 178 predicted SOA (Guo et al., 2014; Fu and Chen, 2017; Yang et al., 2016) (Pankow et al., 2015; Jathar et al., 2016). The secondary portion has been proved to be dominated in PM<sub>2.5</sub> and 179 organic aerosol increases during haze events (Huang et al., 2014;Sun et al., 2019). In 180 181 addition.

182 China has been suffering from severe PM<sub>2.5</sub> pollution especially in the eastern region 183 with fast urbanization and economic development (Fu and Chen, 2017). SOA is a very important component of PM<sub>2.5</sub> in China that contributes about 20-50% (Li et al., 2017b). 184 185 The fraction of SOA in OA increases during haze events (Huang et al., 2014; Sun et al., 2019). Previous modeling studies indicated in China indicate that SOA was underpredicted 186 187 in this region (Wang et al., 2011;Lin et al., 2016;Jiang et al., 2012)(Lin et al., 2016; Jiang et al., 2012) and the impacts of the non-ideality and water-OPM partitioning on modeled 188 189 SOA have not been evaluated.

190 In this study, regional simulations of SOA during January and July of 2013 over 191 eastern China under several scenarios-were conducted to investigate the seasonal variation 192 of SOA due to water partitioning into OPM. Model performances were firstly The model 193 performance was evaluated against observed meteorological parameters (temperature and relative humidity, RH) as well as PM<sub>2.5</sub>, organic carbon (OC<sub> $\tau$ </sub>), and OA at ground 194 195 monitoring sites. Then, the The regional and seasonal impacts on SOA, ALW<sub>org</sub>, and water contentproperties of aerosols were quantified. Factors related to the impacts on SOA, Lastly, 196 197 influences of the results by several factors including sources of precursors, chemical compositions and meteorological conditions were further parameters, estimations of 198 saturation vapor pressures (SVP) of condensed organics, and the individual impacts of 199 200 ALW<sub>org</sub> and non-ideality of the organic-water mixture on SOA prediction were analyzed. Lastly, the impacts on aerosol optical properties and hygroscopicity were investigated. 201

202 2 Methodology

### 203 2.1 Model description

204 The Community Multi-scale Air Quality model (CMAQ v5.0.1) coupled with a modified SAPRC-11 was used in this study. Model configurations were largely based on that used 205 206 by Hu et al. (2016) as summarized below. Firstly, SAPRC-11 was expanded for a more detailed treatment of isoprene oxidation and tracking dicarbonyl products (glyoxal and 207 methylglyoxal) products from different groups of major precursors (Ying et al., 2015); 208 Secondly, heterogeneous formation of secondary nitrate and sulfate from NO2 and SO2 209 reaction on particle surface (Ying et al., 2014), and (Ying et al., 2015). Secondly, SOA from 210 isoprene epoxydiols (IEPOX), methacrylic acid epoxide (MAE) and dicarbonyls through 211 surface-controlled irreversible reactive uptake were added (Li et al., 2015; Pankow et al., 212 213 2015)(Hu et al., 2017; Li et al., 2015; Liu et al., 2020; Ying et al., 2015) were added; Thirdly. Thirdly, the heterogeneous formation of secondary nitrate and sulfate from NO2 214 215 and SO<sub>2</sub> reaction on the particle surfaces (Ying et al., 2014) were added, which is an important source of secondary inorganic aerosols (Zheng et al., 2015) and improves model 216

estimates of nitrate and sulfate (Qiao et al., 2018; Shi et al., 2017). Fourthly, SOA yields
were corrected for vapor wall loss (Zhang et al., 2014)(Zhang et al., 2014)-. Impacts of the
above updates on model performances have been extensively discussed in the cited work
and will not be further investigated in the current study.

221 Two types of SOA were considered in the current model, "semi-volatile" (SV) portion that formed via equilibrium absorption-partitioning of SVOCs, and "non-volatile" (NV) 222 223 portion that formed via direct oxidation of aromatics at low  $NO_{x}$ , isoprene oxidation under 224 acidic conditions, reactive uptake of dicarbonyls, IEPOX and MAE, and oligomers. The 225 SV-SOA module mostly based on that of Pankow et al. The SOA module mostly follows 226 Pankow et al. (2015) (2015) with several updates in the treatment of primary organic aerosol (POA) by including it in the non ideality calculation of the organic water mixture. The 227 mass distribution of SVOCs between the gas phase and particle phase follows the equation: 228 . Two types of SOA as traditionally treated in CMAQ were considered, "semi-volatile" 229 (SV) portion that formed via equilibrium absorption-partitioning of SVOCs, and "non-230 volatile" (NV) portion that includes the oligomers and SOA formed via direct oxidation of 231 aromatics at low-NO<sub>x</sub>. SOA from dicarbonyls, IEPOX, and MAE were formed by 232 irreversible reactive uptake and categorized as NV-SOA in the current model as well. Some 233 studies investigated SOA from glyoxal, methylglyoxal, and IEPOX using detailed 234 reactions and reversible pathways in models or observed as reversible processes in chamber 235 experiments, leading to a relatively lower SOA yield compared to the surface-controlled 236 irreversible uptake (Lim et al., 2013; Knote et al., 2014; Galloway et al., 2009; El-Sayed et 237 al., 2018; Budisulistiorini et al., 2017). The non-volatile assumption used in this paper 238 allows an upper-limit estimation of the importance of these additional SOA formation 239 pathways. POA was treated as non-volatile and non-reactive. The mass distribution of 240 SVOCs between the gas-phase and particle-phase follows the equation: 241

$$K_{p,i} = \frac{F_i}{M \cdot A_i} \tag{Eq 1}$$

where  $K_{p,i}(m^3 \mu g^{-1})$  is the gas/\_particle partitioning constant for compound i,  $F_i(\mu g m^{-3})$  is the concentration of species i in the particle-\_phase,  $A_i((\mu g m^{-3}))$  is the concentration of species i in the gas-\_phase, and  $M_(\mu g m^{-3})$  is the total mass concentration of the absorbing <u>organic</u> phase-\_(i.e. OPM). The gas/\_particle partitioning constant  $K_{p,i}$  is dependent on the <u>chemical</u> composition of the absorbing <u>organic</u> phase.<u>OPM</u>. Pankow et al. (1994)(1994) derived  $K_{p,i}$  for SVOCs partitioning into an absorbing organic phase as:

$$K_{p,i} = \frac{RT}{10^6 \overline{MW} \xi_i p_{L,i}^o} \tag{Eq 2}$$

where  $p_{L,i}^{o}(\text{atm})$  is the saturation vapor pressure<u>SVP</u> of the pure compound i at temperature T(K),  $\xi_i$  is the activity coefficient of species i in the absorbing organic phase,  $\overline{MW}$ (g mol<sup>-1</sup>) is the average molecular weight of the absorbing organic phase<u>OPM</u>, R (8.314 J mol<sup>-1</sup> K<sup>-1</sup>) is the gas constant, and 10<sup>6</sup> is used to convert the <u>unitunits</u> to  $m^3 \mu g^{-1}$ .

There are 12 lumped SVOCs generated by oxidation of alkanes, alkenes, and 253 254 aromatics oxidized under different NO<sub>x</sub> conditions (and 8 NV organic products as listed in 255 Table S1). and Table S2. More details about the lumped precursors such as formation conditions ("high" and "low" NO<sub>x</sub>), lumping species and method, and yields from parent 256 VOCs can be found in Carlton et al. (2010) and summarized in SI. Activity coefficients of 257 SVOCs were calculated based on the composition of absorbing organic phaseOPM using 258 the UNIversal Functional Activity Coefficient (UNIFAC) method (Fredenslund et al., 259 260 1975, method, with assigned carbon number (n<sub>c</sub>), functional groups and energy interaction parameters to both SV and NV compounds (Pankow et al., 2015)(Pankow et al., 2015). 261 The UNIFAC model is one of the commonly used models that activity coefficients of 262 condensed organics and their interactions with water can be estimated. This method has 263 been adopted to investigate the impacts of non-ideality and water-OPM partitioning into 264 OPM on SOA for different precursors in box models (Seinfeld et al., 2001;Bowman and 265 Melton, 2004) (Seinfeld et al., 2001; Bowman and Melton, 2004) and CTMs (Jathar et al., 266

267 2016;Pankow et al., 2015;Kim et al., 2019)(Pankow et al., 2015; Kim et al., 2019). The 268 primary organic aerosols (In the current model, POA) was assumed to have a bulk 269 composition of ten categories of surrogate species (Table S3), as used by Li et al. 270 (2015)(2015). POA is also involved in the calculation of activity coefficients for the 271 organics in the condensed phase.organic-water mixture. Detailed information about the 272 surrogate species including thetheir structures and properties can be found in Li et al. 273 (2015)(2015) and references therein.

In addition to organic compounds, water partitioning into OPM is enabled according to Eq 1 and Eq 2. In such a case, the absorbing phase in Eq 1 includes both organic aerosols and water <u>partitioning intoassociated with</u> OPM. As water <u>consideredcondenses</u> in the absorbing <u>organic</u> phase, it will further alter the molar fraction of each composition, the activity coefficient of SVOCs and the SV-SOA mass concentrations as a result. <u>In the</u> <u>current model, we assumed no interactions between the inorganic and organic phases.</u>

280 As the water partitioning into OPM is highly correlated with the hygroscopicity of 281 aerosols ( $\kappa$ ), their correlation can be expressed by the  $\kappa$ -Köhler theory with Kelvin effect 282 neglected (Peter et al., 2006):

#### 283 **2.2 Estimating the hygroscopicity of OA**

284 <u>Based on the  $\kappa$ -Köhler theory with linearly additive hygroscopic behavior of each</u> 285 <u>component of the mixed particle, ALW<sub>org</sub> is related to the hygroscopicity parameter for the</u> 286 <u>organic mixture ( $\kappa_{org}$ ) by Eq3 (Petters and Kreidenweis, 2007):</u>

$$ALW_{org} = V_{org} = \rho_w V_{org} \kappa_{org} \frac{a_w}{1 - a_w}$$
(Eq3)

where  $\rho_w$  is the density of water (assumed to be 1 g cm<sup>-1</sup>),  $V_{org}$  is the volume concentration of organicorganics, and  $a_w$  is the water activity (assumed to be the same as RH). TakenSince ALW<sub>org</sub>

290 <u>in this study is calculated mechanistically using the partitioning theory,  $\kappa_{org}$  can be 291 <u>estimated by rearranging Eq3:</u></u>

$$\kappa_{org} = \frac{ALW_{org}}{\rho_w V_{org}} \times \frac{1 - a_w}{a_w}$$
(Eq4)

Vora can be estimated from the modeled mass concentration of OA, assuming the density 292 293 of organic aerosolOA to be 1.2 g cm<sup>-3</sup> (Li et al., 2019), the hygroscopicity of the total OA can estimated. This simplified method can be used to estimate OPM associated water. 294 In many studies,  $\kappa_{org}$  is assumed to increase linearly with the oxidation state of OA, 295 296 expressed as the O:C ratio (Guo et al., 2015;Li et al., 2019)(Massoli et al., 2010;Duplissy 297 et al., 2011;Lambe et al., 2011). In addition, the hygroscopicity of organic aerosol is dependent on the degree of oxygenation, showing a positive linear relationship with the 298 O:C ratio (Massoli et al., 2010; Duplissy et al., 2011; Lambe et al., 2011; Hong et al., 2018; Li 299 et al., 2019). We therefore estimated the correlation of  $\kappa$  and O:C ratio at 9 representative 300 cities during January and July. The correlation of  $\kappa_{org}$  and O:C ratio at 9 representative 301 cities was evaluated during January and July of 2013, with the reduced major axis 302 regression method (Avers, 2001). The O:C ratio of the total OA was calculated as 303 304 followingusing Eq5:

$$0: C = \sum_{i=1}^{n} f_i (0:C)_i \qquad (\underline{\mathsf{Eq4Eq5}})$$

where  $f_i$  and  $(O:C)_i$  are the molar fraction and O:C ratio of organic aerosol component i. For POA, a fixed molar fraction and composition has beenwere assumed following Li et al. (2015)(2015). For SOA, the O:C ratio was estimated by their OM:OC ratio (Simon and Bhave, 2012):

309 . For SOA, the O:C ratio was calculated by using their organic matter to organic carbon
 310 ratio (OM:OC) following Simon and Bhave (2012):

$$O:C = \frac{42}{45} = \frac{12}{15} (OM:OC) - \frac{14}{15}$$
 (Eq5Eq6)

311 OM:OC ratio of each SOA component follows Pye et al. (2017).

312 The OM:OC ratio of each SOA component follows Pankow et al. (2015) as shown in Table
 313 S1-S2.

# 314 2.3 Model application

315 The simulation domain has a horizontal resolution of 36 km  $\times$  36 km (100  $\times$  100 grids) 316 and a vertical structure of 18 layers up to 21 km, which covers eastern China as shown in Figure S1. Anthropogenic emissions were generated from the Multi-resolution Emission 317 318 Inventory for China (MEIC) (Zhang et al., 2009:Li et al., 2014:Zheng et al., 2014:Liu et al., 2015)(Zhang et al., 2009) v1.0 with a  $0.25^{\circ} \times$ 319  $0.25^{\circ}$ resolution 320 (http://www.meicmodel.org) for China, and the Regional Emission inventory in Asia version 2 (REAS2) (Kurokawa et al., 2013) with a  $0.25^{\circ} \times 0.25^{\circ}$  resolution 321 (http://www.nies.go.jp/REAS/) for the rest of the domain. Biogenic emissions were 322 323 generated by the Model for Emissions of Gases and Aerosols from Nature (MEGAN) v2.1, 324 with the leaf area index (LAI) from the 8---day Moderate Resolution Imaging Spectroradiometer (MODIS) LAI product (MOD15A2) and the plant function types (PFTs) 325 from the Global Community Land Model (CLM 3.0). Open biomass burning emissions 326 327 were generated from the Fire INvnetory from NCAR (FINN) (Wiedinmyer et al., 328 2011)(Wiedinmyer et al., 2011). Dust and sea salt emissions were generated in lineonline during CMAQ simulations. The total emissions of major SOA precursors and their spatial 329 330 distributions are shown in Table S4 and Figure S2. Meteorological fields were generated using the Weather Research and Forecasting (WRF) model v3.6.1 with initial and boundary 331 conditions from the NCEP FNL Operational Model Global Tropospheric Analyses dataset. 332 More details about the model application can be found in Hu et al. (2016) 333

Four scenarios are investigated in this study. The base case (BS) that applied the default secondary organic aerosol module of CMAQ; the water case (S1) that only water partitioning into OPM was considered; the UNIFAC case (S2) that effects of molecular structure of the primary and secondary organic species were included; and the combined case (S3) that S1 and S2 were combined together. The base case (BS) applies the default secondary organic aerosol module of CMAQ v5.0.1. In this case, no water partitioning into OPM is considered. Lumped semi-volatile products from the oxidation of various 341 precursors partition into a single organic phase, which is considered as an ideal mixture of POA and SOA with  $\gamma_{org}=1$ . The water case (S1) includes water partitioning into OPM, 342 which is again considered as an ideal solution ( $\gamma_{org}=1$  and  $\gamma_{H2o}=1$ ). The UNIFAC case 343 (S2) considers the interaction between organic constituents with UNIFAC calculated 344 activity coefficients ( $\gamma_{org} \neq 1$ ) but does not allow water partitioning into OPM. The 345 346 combined case (S3) allows both water partitioning and interactions between all constituents 347 (including water and organics) using UNIFAC calculated activity coefficients ( $\gamma_{org} \neq 1$ and  $\gamma_{H20} \neq 1$ ). The results of BS and S3 are used to examine the overall impacts of water 348 partitioning into OPM and polarity of organics on SOA and ALWorg, as shown in Section 349 3.1-3.4. The separate influences of those two processes on SOA from S1 and S2 are 350 351 discussed in Section 3.5. **3 Results and discussion** 352 353 3.1 Model evaluation 354 Temperature The meteorological inputs and relative humidity (emissions have been used in 355 several previous publications. Model performance on meteorological parameters 356 (temperature and RH), gaseous species and gas and aerosol concentrations have been extensively evaluated (Hu et al., 2016; Hu et al., 2017; Qiao et al., 2018; Shi et al., 2017). 357 A summary of the model performance related to this study is provided below. Observed 358 meteorological data were obtained from the National Climatic Data Center 359 (ftp://ftp.ncdc.noaa.gov/pub/data/noaa). Observations of OC at two urban locations, 360 Beijing (Cao et al., 2014; Wang et al., 2015) and Guangzhou (Lai et al., 2016) and OA in 361 362 Beijing (Sun et al., 2014) during January of 2013 as well as surface PM<sub>2.5</sub> at several monitoring sites during July of 2013 from China National Environmental Monitoring 363 364 Center (http://113.108.142.147:20035/emcpublish/) were used to evaluate model estimates of aerosols. Details of measurement methodology and uncertainties of observations are 365 listed in the corresponding references. 366

367 Temperature and RH are the two meteorological factors that affect SOA formation. Table 1 shows the comparison of lists model statistics of mean observation (OBS), mean 368 369 prediction (PRE), mean bias (MB), gross error (GE) and correlation coefficient (R) based 370 on WRF predictions and observations at monitoring sites located in 8 sub-regions of the domain (Figure S1). Observed data are accessible from the National Climatic Data Center 371 at ftp://ftp.ncdc.noaa.gov/pub/data/noaa. Temperature and RH are ) during January and 372 373 July of 2013. The benchmarks for the MM5 model (another meteorology model) of 4-12km horizontal resolution suggested by Emery et al. (2001) are also listed in the table. Details 374 of monitoring sites in the 8 sub-regions are listed in Table S5. Overall, WRF tends to 375 376 underestimate both temperature and RH. The model shows better agreement with observed temperature as R is higher than that of RH. Both temperature and RH are well captured by 377 WRF the model in YRD, the Pearl River Delta (PRD), and the central regions of China (the 378 major regions of eastern China). In these regions, MB and GE of temperature are -1.2~0.7 379 K and 1.8~2.6 K, respectively, which are -11.8~5.6% and 9.2~16.8% for RH, respectively. 380 381 Model estimates of daily organic carbon (OC) from the BS case were S3 are compared with measurements at monitoring sites in Beijing and Guangzhou during the winterin January 382 383 of 2013 (Figure 1(a)). Overall, The factors used to convert SOA to OC (SOC) are listed in Table S1-S2. OC from POA (POC) is directly predicted by the model. Generally, the ratio 384 between modeled and observed OC concentration falls in the range of 1:2 to 2:1, with a 385 correlation coefficient R of 0.70. The model tends to underestimate OC, especially in 386 387 Beijing on highly polluted days (by -37-48%). No significant improvements to modeled OC were observed in S3. The impacts of water co-condensation and polarity of organic 388 condensed species on SOA exhibit strong seasonal and spatial features, which are further 389 discussed in Section 4. The impacts in Beijing and Guangzhou are not significant during 390 winter. an R-value of 0.7. The model tends to underestimate OC on high concentration 391 392 days. Overall, the mean fractional bias (MFB) and mean fractional error (MFE) of OC are -0.20 and 0.27, within the criteria (MFB $\leq \pm 0.6$ ; MFE $\leq 0.75$ ) suggested by EPA (2007). 393

The bias in OC might be due to <u>under-estimated</u> POA emissions and <u>under-predicted</u> POA emissions and <u>under-predicted</u> SOA in CMAQ from missing precursors (Hu et al., 2017;<del>Zhao et</del> al., 2016a); Zhao et al., 2016a).

397 The model estimate of OA was further investigated. As. No significant differences in
 398 OC are observed in S3 compared to BS (not shown-in), likely due to the biased-low SOA
 399 predicted in the current model that limits the impact of ALW<sub>org</sub> on SOA formation.

400 The underestimate of SOA can be seen from Figure  $1(b_{\tau})$  as well. CMAQ can well 401 capture the observed diurnal variation of OA atin Beijing during wintertime, except for the underestimates of peak values. A better agreement The correlation coefficient of modeled 402 to observed OA is 0.55. We find a 25% underestimate of OA on average. Better agreement 403 between the model and the observations is observed shown on non-polluted days (daily-404 averaged concentration less than 75 µg m<sup>-3</sup>). The monthly averaged mean fractional bias 405 (MFB) and mean fractional error (MFE) are -0.13 and 0.27, respectively. POA is The mean 406 fractional bias (MFB) and mean fractional error (MFE) of polluted days are -0.38 and 0.64, 407 respectively, which are worse than that of the non-polluted days (-0.26 for MFB and 0.52 408 409 for MFE). The overall MFB and MFE of OA during January are -0.28 and 0.54, within the 410 criteria (MFB $\leq \pm 0.6$ ; MFE $\leq 0.75$ ) suggested by EPA (2007). Again, no apparent changes of SOA nor OA are observed between case S3 and BS (not shown), since POA is predicted 411 to be the primary contributor to OA at Beijing in winter, accounting for 88% due to aging 412 of POA not treated in the current model. The fraction of SOA is small, resulting in little 413 impacts on SOA by in the current model, with an averaged SOA/POA ratio of 0.12. This 414 ratio is much lower than the field observation of about 0.45-1.94 (Zhao et al., 2019; Sun et 415 al., 2013; Sun et al., 2016). The bias might be due to the missing SOA converted by 416 partitioning and aging of semi-volatile POA as well as oxidation from intermediate volatile 417 organic compounds (IVOCs) and VOC oxidation products. Those pathways are shown to 418 419 be important for SOA formation by modeling, field and chamber studies (Hodzic et al.,

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420 <u>2010; Jimenez et al., 2009; Murphy et al., 2017; Robinson et al., 2007; Shrivastava et al.,</u>
421 <u>2008; Tkacik et al., 2012; Zhao et al., 2014; Zhao et al., 2016a).</u>

A sensitivity test was performed by using the newest CMAQ model version 5.3.1 that includes all the above processes in the aerosol module. The SOA/POA ratio in Beijing is improved greatly to be 0.83 in winter. However, high uncertainties still exist in the emissions of the involved precursors and characterization of SOA formation through these processes, needing further constrains by observations. Their influences on water partitioning into OPM and insignificant improvements<u>non-ideality</u> of the modeled OAorganic-water mixture on SOA will be evaluated in <u>S3-a</u> future study.

Due to the lack of observed OC and OA in July of 2013, as an alternative, model 429 performances are evaluated by comparing predicted and observed PM<sub>2.5</sub> at ground sites 430 (Figure S2 shows the comparison of modeled and observed PM<sub>2.5</sub> at monitoring sites S1) 431 as shown in Figure S1 (a) during July of 2013.S3. Generally, ourthe model can well 432 reproduce the diurnal variation of PM2.5 in most regions. Predicted PM2.5 on high 433 concentration highly concentrated days are biased low compared to observations, 434 especially in the North CentralChina Plain (NCP). The NCP region has the highest PM<sub>2.5</sub> 435 ranging from 60 µg m<sup>-3</sup> to 300 µg m<sup>-3</sup> compared to other regions. The bias in modeled 436 PM<sub>2.5</sub> is significant in cities in the Northwest. This might be due to missing dust emissions 437 in the current inventory (Hu et al., 2016). To further evaluate the model performance, 438 statistics of averaged MFB and MFE were of modeled PM<sub>2.5</sub> are plotted against observed 439 PM<sub>2.5</sub>-concentration at all monitoring sites (observations of each site as shown in Figure 440 S3).S4. The criteria of MFB and goalMFE followed recommendations of Boylan and 441 Russell (2006) (2006). Our model performed performs well assince most of the predictions 442 meet the criteria and a large fraction (>58%) meet the goal. The averaged MFB and MFE 443 444 across all the sites are -0.28 and 0.39 respectively, indicating slightly underestimate of 445  $PM_{2.5}$  by the model.

446 3.2 Impacts of water partitioning on SOA and ALW org

447 Distribution The spatial distribution of SOA varied varies greatly in the two seasons. In 448 winter, SOA is relatively high in eastern SCB and in the contiguous areascentral and eastern 449 provinces of Shandong, Henan, Anhui, and Hubei provinces (Figure 2 and Figure S4S5). 450 Monthly-averaged SOA concentrations in the above two-areas are up to 25 and 15-20 µg m<sup>-3</sup>, respectively. The major precursors of SOA Anthropogenic emissions are originated 451 from anthropogenic the major sources of SOA (Figure S6), such as dicarbonyl products of 452 453 aromatics from the oxidation, xylenes of xylene and toluene (Figure S5). (Hu et al., 2017). In summer, surface SOA is high in **NENortheast**, NCP, and YRD regions. Shanghai, 454 Jiangsu province and coastal areas of Yellow Sea show the. The highest SOA of ~9 12 µg 455 m<sup>-3</sup>-occurs in Shanghai and Jiangsu provinces as well as the coastal area of the Yellow Sea, 456 with the value of ~9-16  $\mu$ g m<sup>-3</sup> at the surface and ~20-25 mg m<sup>-2</sup> as in the column total (col-457 SOA) inof the atmosphere below 21 km (Figure <u>\$4\$5</u>). Different from winter SOA, a 458 significant fraction of summer SOA is originated from biogenic emissions in Shanghai and 459 Jiangsu province (Figure S5S7). Anthropogenic SOA is high-highly concentrated in July 460 461 in the coastal areas of the Yellow Sea and Bohai Bay.

Combined effects of water partitioning into OPM and non-ideality on SOA formation 462 (S3)-also exhibit strong seasonal variation. In winter, the increase of SOA is relatively 463 small, slightly decreased by  $-1-4.5 \ \mu g \ m^{-3}$  (10-20%) on average at the surface (Figure 2) 464 and less than  $\sim 51 \text{ mg m}^{-2}$  (10.30%) as for 20%) in the column concentration (Figure S4). 465 The influences on SOA also differ in different altitudes. For example, the maximum 466 increment at the surface is observed in Shandong province in NCP (Figure 2), while SOA 467 at higher levels of the atmosphere is more significant in South China (Figure S4). The 468 increase in SOA is mostly attributed to anthropogenic sources in winter (Figure S5-and 469  $\frac{S7}{2}$ ) over high SOA regions where anthropogenic sources dominate. We show later that 470 the decrease of SOA is mainly due to the large activity coefficients of SVOCs which 471 472 <u>decrease</u>  $K_{p,i}$ . In summer, higher temperature and <del>relative humidity (RH)</del> promote <u>water</u> 473 partitioning and SOA formation as well as water partitioning into OPM. At the surface, so 474 that SOA increases by 3-9 µg m<sup>-3</sup> (40-50%) in coastal areas and 2-9 µg m<sup>-3</sup> (20-90%) over the land, which are dominated by anthropogenic and biogenic origin, respectively (Figure 475 476 S6). For col-SOA, in addition to coastal areas, more significant increase is observed in 477 YRD, most of Henan province, and the contiguous areas of Hubei, Hunan, and Jiangxi province (Figure S4) by about 30-70% apparently over the entire domain, with the highest 478 enhancement of 2-4  $\mu$ g m<sup>-3</sup> (20-50%) at the surface (Figure 2) and 4-6 mg m<sup>-2</sup> (30-60%) in 479 480 the column (Figure S5) over YRD and the coastal area of Yellow Sea. Anthropogenic SOA dominates the total change in winter as shown in Figure S6. In summer, the increase of 481 SOA is attributed to biogenic sources in inland areas and anthropogenic sources over the 482 ocean (Figure S7). 483

Regional distribution of water partitioning into OPM ALWorg is similar to the 484 changeschange of SOA- as shown in Figure 3-shows the regional distribution of monthly-. 485 In winter, a maximum averaged daily maximum ALW<sub>org</sub>. We see up to 9 concentration of 486 3.0 µg m<sup>-3</sup> for ALW<sub>org</sub> at surface occurs in Shandong in winter where great increment inthe 487 high SOA appears as well. In region, where significant changes of SOA also occur. In other 488 areas, the averaged concentration of ALW<sub>org</sub> is about  $2-60.5-1.5 \mu \text{g m}^{-3}$ . TheOverall, the 489 average ratio of ALW<sub>org</sub> to SOA is about 0.1-0.53 in winter. In summer, water partitioning 490 into OPM mostly involvesoccurs in the east coastal areasarea at the surface where a 491 492 significant increase of anthropogenic SOA (such as those from toluene and xylenes) is observed. This might be due to the high polarity of anthropogenic SVOCs (having more -493 COOH groups) that absorb more water. In the coastal areas, ALW<sub>org</sub> is about 5-12  $\mu$ g m<sup>3</sup>, 494 with a ratio to SOA area, the averaged concentration of 0.3-0.6. ALW org over the land is 495 about 25-7 µg m<sup>-3</sup>, with the ALW<sub>org</sub> /SOA ratio of 0.5-1.0. Over the land, the averaged 496 concentration of ALWorg is about 1-3 µg m<sup>-3</sup> (ALWorg/SOA ratio of 0.+2-0.45) in most 497 areas, which Northeast and East China. Water partitioning is mostly associated with the 498 499 increase of BSOA such as originated from isoprene and monoterpenes oxidation that 500 produces SVOCs with abundant OH group in SVOCs. The highest.

501 Based on the column concentrations of ALWorg is 16 µg m<sup>-3</sup> near Shanghai (and ALW<sub>org</sub>/SOA ratio of 0.57). Water partitioning also varies at different altitudes (Figure S9). 502 In (Figure S8), in winter, more column water partitions into OPM (col-ALWorg) in 503 Chongqing, Hunan, Guanxi, Guangdong and Guizhou province, with the ALW org must have 504 occurred in the south and southwest regions at higher levels where a significant increase of 505 col-SOA occurs (Figure S5). The averaged col-ALW<sub>org</sub>/col-SOA ratio of in the high SOA 506 507 area is 0.21-0.3. In summer, the ALW<sub>org</sub> must be high at higher altitudes over the central regions in China. The maximum col-ALWorg is predicted about 7 mg m<sup>-2</sup> over the land, 508 especially in YRD, with the col-ALW<sub>org</sub>/col-SOA ratio of 0.1-about 0.3-over eastern China. 509 510 Figure 4 shows 3.3 Impacts on aerosol properties Since ALW<sub>org</sub> is determined in S3, the correlation of  $\kappa_{org}$  with O:C ratio. The values 511 can be estimated from the modeled ALW<sub>org</sub>, OA and RH using Eq 4. 9 representative cities 512 513 were selected to investigate  $\kappa_{org}$  vs. O:C relationship and its seasonal variation as shown in Figure S9 and S10. The results of all the cities in winter and summer are merged and 514 analyzed as shown in Figure 4. Pairs of  $\kappa_{org}$  and O:C data are grouped into 10 O:C bins 515 and the average  $\kappa_{org}$  in each bin are then calculated. Overall, the estimated O:C ratio is 516 517 within the range of 0.2-0.6. In summer, the oxidation state of OA shows different degrees of enhancement compared to winter at most of the cities except Guangzhou, due to 518 increased contribution of SOA to total OA. The averaged  $\kappa_{org}$  of OA in each O:C bin 519 falls in the range of 0.001-0.18. The averaged  $\kappa_{org}$  in each O:C bin is less than 0.1 in 520 winter, with the highest  $\kappa_{org}$  (~0.3) at Beijingvalue in Guangzhou. As more ALW<sub>org</sub> is 521 formed in summer-, the averaged  $\kappa_{org}$  also increases greatly with the highest value of 522 <u>0.35 in Beijing</u>. The linear correlation between  $\kappa_{org}$  and O:C shows significant spatial 523 and seasonal variations. For example, the slope of  $\kappa_{org}$ -O:C is <u>much70-90%</u> smaller in 524 winter (45-74% less) than in summer in the Northern cities such as Shenyang, Beijing, 525 Zhengzhou, and Xi'an, while. However, in Guangzhou, the slope of  $\kappa_{org}$ -O:C in winter is 526 muchis 83% higher (47-104% more) in winter than in summer in the Southern cities, such 527

528 as Nanjing, In Chengdu and Guangzhou. In Jinan and Shanghai, the slope is quite similar in both seasons. The fitted correlations are very different from previous studies with a 529 relatively higher Overall, the slope of  $\kappa_{org}$ -O:C from vs. O:C at the 9 cities is 0.16 in winter 530 531 and 0.40 in summer. Most of the fitted linear correlations of the individual city fall outside of the range of 0.18 to 0.37 suggested in previous studies (Duplissy et al., 2011;Lambe et 532 al., 2011; Massoli et al., 2010; Chang et al., 2010) (Duplissy et al., 2011; Lambe et al., 533 534 2011; Massoli et al., 2010; Chang et al., 2010), indicating that the hygroscopicity of organic 535 aerosols with chemical complexity cannot be simply be represented by a single parameter such as the O:C ratio (Rickards et al., 2013)(Rickards et al., 2013).-536

## 537 **3.3 Impacts on solar radiation**

538 <u>. In both seasons,  $\kappa_{org}$  approaches zero and negative values as O:C decreases, which</u> 539 <u>might be due to the linear regression of  $\kappa_{org}$  and O:C. To avoid this, an exponential fitting</u> 540 <u>of the two variables is performed so that  $\kappa_{org}$  falls in the range of (0,1) and is positively</u> 541 <u>correlated with O:C. In this case, the fitted correlations are  $\kappa_{org}=1-\exp(-(O:C/1.88)^{2.29})$ </u> 542 <u>and  $\kappa_{org}=1-\exp(-(O:C/1.06)^{4.50})$  for January and July of 2013, respectively..</u>

The impacts on aerosol optical depth (AOD) and aerosol radiative forcing (ARF) wereare further investigated. Figure 5 shows the monthly-averaged AOD at 550 nm in January and July of 2013. It wasis calculated asby summarizing the accumulationproduct of model estimated extinction coefficient of fine particles ( $EXT_ib_{ext,i}$ ) multiplied by the thickness (*HLi*) of an each layer:

$$AOD = \sum_{i=1}^{N} EXT_{i} \sum_{i=1}^{N} b_{ext,i} \times HL_{i}$$
(Eq6)

548 Wherewhere N is the number of layers. There are two methods to estimate the aerosol 549 extinction coefficient in CMAQCMAQv5.0.1. One is usingbased on the Mie theory 550 (*EXT*<sub>m</sub>and the predicted aerosol component concentrations ( $b_{ext,m}$ ), and the other is based 551 on extinction values<u>correlation</u> from the IMPROVE monitoring network that considers the 552 impacts of hygroscopicity of different aerosol components (*EXT*<sub>r</sub>)<u>b\_{ext,r</u>) (Malm et al., 553 <u>1994)(Malm et al., 1994)</u>. AOD calculated with the two types of extinction coefficient are
 554 denoted as AOD<sub>m</sub> and AOD<sub>r</sub>, respectively.

555 In Figure 5, a clear pattern of high AOD<sub>r</sub> in SCB and NCP and low AOD<sub>r</sub> in west 556 China is observed in both winter and summer is observed, consistent with previous studies (He et al., 2019; He et al., 2016; Luo et al., 2014)(He et al., 2019; He et al., 2016; Luo et al., 557 558 2014). An identified identical trend in AOD<sub>m</sub> is observed as shown in Figure  $\frac{$10}{$10}$  S11. The monthly-averaged AOD<sub>r</sub> ranges from 1.40 to 3.52 in January and from 0.43 to 0.89 in July. 559 560  $AOD_m$  is lower than  $AOD_r$ , falling in 0.7-2.2 in January and 0.3-0.6 in July. The model significantly overestimates AOD in January but agrees better with observations from 561 562 MODIS in the where AOD is high regions in July (Figure S11). Biases S12). The bias in the 563 predicted AOD might be partially due to the empirical equation applied in the calculation 564 of AOD in CMAQ (Wang et al., 2009;Liu et al., 2010) (Wang et al., 2009; Liu et al., 2010), and partially due to the uncertainties of fine AOD overland from MODIS data (Wang et al., 565 2009;Levy et al., 2010)(Wang et al., 2009; Levy et al., 2010). With water partitioning into 566 567 OPM, changes in SOA mass concentration and chemical composition lead to The increase of AOD, which due to ALWorg shows a strong spatial and seasonal pattern. In winter, there 568 569 isare no significant increase changes in AOD<sub>r</sub> across the whole domain, due to insignificant changes of SOA. In summer, AODr increases significantly in YRD and the adjacent area 570 of Hubei, Hunan, and Jiangxi province areas by up to 1210%. 571

ARF represents the changeschange in the radiative flux at the top of the atmosphere 572 573 due to aerosols. The off line An offline version of the Shortwave Radiative Transfer Model 574 For GCMs (RRTMG\_SW) is was used to calculate the direct radiative effect of aerosols on shortwave radiation (Iacono et al., 2008)(Iacono et al., 2008). Generally, fine aerosols 575 576 exhibit cooling effects on the shortwave radiation in both winter and summer over the entire 577 domain as shown in Figure 6. This impact is much stronger in the areas where AOD is high 578 (Figure 5). The ARF at top of atmosphere (TOA) is highest in Shandong in winter and in the coastal areasarea near Jiangsu province in summer, which are about  $-\frac{125}{125}$  W m<sup>-2</sup> and -579

96 W m<sup>-2</sup>, respectively. In winter, no significant changes of ARF are observed in the high regions of eastern China (Figure 66b and d). This is likely attributed to an insignificant contribution of SOA to PM<sub>2.5</sub> in winter compared to other components with cooling effects, such as sulfate. In summer, SOA is an important component of PM<sub>2.5</sub> (20-60%), and the effects of water partitioning on shortwave radiation isare relatively stronger. An enhancement of up to 15% in the cooling effects of ARF occurs near the YRD region where AOD significantly changes as well.

587 **4 Discussion** 

### 588 3.4 Sensitivity to T, RH, and SVP

589 Meteorological conditions and SOA precursors affect the impacts of water partitioning on 590 SOA. Figure 7 shows the effects of different factors on the daily maximum change of SOA in Jinan and Nanjing, two representative cities in winter and summer, respectively. As 591 shown in Figure 7(a), the daily maximum elevation of SOA occurs when RH is greater than 592 70% in both cities. This is consistent with the previous study in the Southeast U.S. during 593 594 summer (Pankow et al., 2015). A clear correlation of the changes in SOA with SOA concentration in Nanjing (R=0.84) during summer can be observed. However, this 595 correlation is relatively weak in Jinan (R=0.44) during winter. There is no strong 596 correlation between changes in SOA and temperature as shown in Figure 7(b), likely due 597 to the daily variation of SOA mass and composition. To better illustrate the dependency of 598 SOA on temperature and, relative humidity, and SVP of SVOCs, an offline calculation of 599 600 SOA formation was performed at two representative cities (Jinan in NCP during winter 601 and Nanjing in YRD during summer) when the daily maximum SOA increases occurred. We assumed assume temperature (T) and water vapor mixing ratio (QV) to be within the 602 603 range of  $\overline{X} \pm \sigma$ , where  $\overline{X}$  and  $\sigma$  are the mean and standard deviation calculated based on 604 WRF predictionpredictions at each location. We chose choose 10 evenly distributed values 605 for T and QV within the range of  $\overline{X} \pm \sigma$ . The temperature dependence parameter of 606 saturation vapor pressure SVP ( $\Delta H$ ) was/R) is also scaled by 0.2, 0.8, 1.4 and 2.0 separately 607 for all the SVOCs. As shown in Figure 87, SOA indicates exhibits a negative correlation with temperatureT and a positive correlation with RHOV in both cities. SOA is more 608 sensitive to RHQV under coolcold conditions (JNJinan) and to temperature under hot 609 610 conditions (NJ). An interesting finding is that significant increases in SOA Nanjing). When the temperature is fixed, the sensitivity of SOA to  $\Delta H/R$  is different in the two cities occur 611 during different time periods of the day. Water partitioning tends to affect SOA in the 612 613 afternoon and evening in Jinan, which mostly happens in the early morning and at noon in Nanjing. The different timing is likely. We find more changes in SOA across  $\Delta H/R$  in 614 Jinan. This is attributed to a substantial increase in SOA precursors in the two cities. In 615 Jinan, the most contributing SVOCs are originated the temperature correction factor ( $\zeta_{corr}$ ) 616 of  $K_p$  in CMAQ as defined below: 617

$$\zeta_{corr} = \frac{K_{p,T_{ref}}}{K_{p,T}} = \frac{T_{ref}}{T} \exp\left[\frac{\Delta H}{R} \left(\frac{1}{T_{ref}} - \frac{1}{T}\right)\right]$$
(Eq8)

where  $T_{ref}$  is the reference temperature (298K) and T is temperature. According to 618 Figure 7, the range of T is 265-274K in Jinan and 300-307K in Nanjing. The deviation of 619 temperature from toluene and xylenes oxidation, as well as oligomers formed by their 620 oxidation products in OPM. Possible emission sources include transportation, petroleum 621 refining, manufacturing, painting, etc. SOA increase in Nanjing is mostly associated with 622 biogenic sources including isoprene and monoterpenes. the reference value (298K) is 623 greater in Jinan than in Nanjing. Therefore, the unit change of  $\Delta H/R$  causes greater 624 variations of  $\zeta_{corr}$  and thus  $K_p$  in Jinan. As a result, SOA is more sensitive to  $\Delta H/R$ . 625 The impacts of SVP estimation on SOA are more significant in winter than in summer. 626 3.5 Separate impacts of water partitioning/polarity of condensed organics 627 Impacts of water partitioning into OPM and non-ideality of the organic-water mixture on 628 SOA are <u>in opposite directions</u>. Water partitioning alone increases SOA by  $\sim 10-20-60\%$  in 629

630 winter and  $\sim \frac{20 - 10030 - 80}{\%}$  in summer <u>in most areas of the domain (Figure S128</u>). This is

631 because-that the molecular weight of water is quite small and will reduce the molar

632 averaged weight of OPM ( $\overline{MW}$ ) in Eq 2 (Pankow et al., 2015)(Pankow et al., 2015). The reduced  $\overline{MW}$  further increases  $K_{p,i}$  promoting the  $K_p$  and promotes mass transfer of 633 SVOCs from the gas-\_phase to the OPM. On the other hand, by considering non-ideality of 634 635 the organic-water mixture, activity coefficients of SVOCs are usually greater than 1.0 in this study, leading to a decrease in  $\frac{K_{p,t}}{K_p}$ . As a result, the total SOA concentration is 636 637 reduced by  $\frac{10}{100} \sim 10^{-20\%}$  in winter and  $\sim \frac{30}{100} \sim 10^{-50\%}$  in summer in the high regions (Figure 638 \$139). Overall, the final impacts are the combined consequences of the two "processes". 639 In winter, the increase of SOA caused by water partitioning is offset by the decrease of 640 SOA due to the polarity of SVOCs in most areas of the domain, resulting in no significant changes.slight decreases of SOA. In summer, effects the effect of water partitioning 641 642 overcome overcomes that of SVOC polarity so as the total SOA loading increases. This further leads to an enhanced attenuation of shortwave solar radiation and cooling of the 643 644 atmosphere.

645 **5 Conclusion** 

646 <u>4 Conclusions</u>

647 The WRF/CMAQ model was used to investigate the impacts of water partitioning into 648 OPM and non-ideality of the organic-water mixture on SOA formation over eastern China 649 during January and July of 2013. SOA is greatly enhanced in summer especially in YRD 650 and over the Yellow Sea by up to 9050% and 7060% at the surface and in the whole column, respectively. No significant impacts on SOA are observed in winter. This might be due to 651 652 the underestimation of SOA in the current model. ALW<sub>org</sub> is highly correlated with the changeschange of SOA, with the ratio of ALWorg to SOA of 0.1-0.53 and 0.2-1-0.0-6 at the 653 654 surface where significant changes of SOA occur in winter and summer, respectively. By 655 using the modeled ALW<sub>org</sub>, correlations between  $\kappa_{org}$  and O:C wereare examined in 9 representative cities, showing significant spatial and seasonal variations. The increases in 656 SOA lead to 12% elevation of AOD and 15% an enhancement in the averaged AOD and 657 658 the cooling effects of ARFaerosols, by up to 10% and 15% respectively in summer. The

659 model predicted SOA is sensitive to temperature and QV in both seasons, with higher sensitivity to QV during winter and temperature during summer. Estimation of SVP also 660 affects modeled SOA, especially in a cold environment. The effects of water partitioning 661 into OPM and non-ideality of the organic-water mixture on SOA were also examined 662 separately are the opposite. Since the activity coefficients of SVOCs are mostly greater than 663 1.0 during the simulated episode, SOA concentrations decrease concentration decreases 664 665 when the non-ideality effect is considered. DailyAveraged SOA concentration decreases 666 by up to -1020% in winter and -3050% in summer in the high regions. Water partitioning alone increases SOA by  $\sim 10-20-60\%$  in winter and  $\sim 20-10030-80\%$  in summer. It should 667 be noticed that the results shown in this study are the lower limit as the current model tends 668 669 to underestimate SOA. It is crucial to consider both effects in simulating SOA formation under hot and humid conditions in CTMs. 670

671

### 672 Data availability

Data used in this manuscript can be provided upon request by e-mail to the corresponding
authors, Qi Ying (qying@civil.tamu.edu), and Jianlin Hu (jianlinhu@nuist.edu.cn).

675

#### 676 **Competing interests**

- 677 The authors declare that they have no conflict of interest.
- 678

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Table 1. Statistical analysis of modeled temperature (K) and relative humidity (%) of January and July of 2013 at the monitoring sites in different geographical areas as shown in Figure S1- and listed cities in Table S5.

\*\*Northeast is Northeast China, NCP is North China Plain, Northwest is Northwest China, YRD is Yangtze River Delta, Central is Central China, SCB is Sichuan \*OBS is mean observation, PRE is mean prediction, MB is mean bias, GE is gross error, R is the correlation coefficient of predictions vs. observations.  $\underline{\text{MB}} = \sum_{i=1}^{N} (C_p - C_o) / N; \quad \underline{\text{GE}} = \sum_{i=1}^{N} |C_p - C_o| / N, \quad \underline{\text{where } C_p \text{ and } C_o \text{ are the prediction and observation data, N is the total number of data.}$ 

Basin, PRD is Pearl River Delta, Southwest is Southwest China.

\*\*\*Benchmarks of MB and GE are for the MM5 model of 4-12 km horizontal resolution by Emery et al. (2001)


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**Figure 1**. Comparison of (a) observed and modeled organic carbon concentration at University of Beihang (BH), Tsinghua University (TH) and Guangzhou (GZ); (b) observed organic aerosol (Obs-)-OA) at Beijing and predictions of total OA (pOA) and SOA (pSOA), unit is µg m<sup>-3</sup>. Locations of monitoring sites are shown in Figure S1.



Figure 2. Monthly-averaged total SOA in BSfrom S3 and monthly-averaged daily

maximum changes of SOA due to water partitioning <u>into OPM</u> and non-ideality of <u>the</u> organic-water mixture. "Abs. Diff." represents absolute differences (S3-BS); "Rel. Diff." represents relative differences ((S3-BS)/BS, %). Relative differences are shown in areas with monthly-averaged SOA concentration greater than  $1 \mu g m^3$ .





**Figure 3**. Monthly-averaged daily maximum water partitioning into the organic-phase (ALW<sub>org</sub>,  $\mu$ g m<sup>-3</sup>) and the ratio to SOA (ALW<sub>org</sub>/SOA) during January and July of 2013. AWL<sub>org</sub>/SOA is shown in areas with monthly-averaged SOA concentration greater than 1  $\mu$ g m<sup>-3</sup>.





**Figure 4.** The correlation of hygroscopicity of <u>organic</u> aerosol ( $\kappa$ ) $\kappa_{org}$ ) and O:C ratio at 9 representative cities including Shenyang (SS), Beijing (BJ), Jinan (JN), Zhengzhou (ZZ), Xi'an (XA), Nanjing (NJ), Shanghai (SH), Chengdu (CD), and Guangzhou (GZ). Gray dots on the background represent all the data) in January (a) and July, which (b) of 2013. O:C ratios are categorized into several O:C10 bins. In each bin, the ranges of  $\kappa$ O:C and O:C ratio $\kappa_{org}$  are represented by dashed bars colored for January (navy) and July (light blue), with the mean value. The mean values of O:C and  $\kappa_{org}$  are represented by triangles colored by the averaged RH of each bin. The mean  $\kappa$ The relationship between  $\kappa_{org}$  and O:C ratio areis fitted by a linear function with reduced major axis regression.— (blue lines) and an exponential function (red lines), respectively.  $\kappa_{01}$  and  $\kappa_{07}$  represent the fitted correlation for January and July, respectively.

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**Figure 5**. Monthly-averaged AOD<sub>r</sub> at 550 nm and the monthly-averaged daily maximum changes of AOD<sub>r</sub> due to water partitioning into OPM and the non-ideality of the organic-

water mixture. "Abs. Diff." represents absolute differences (S3-BS); "Rel. Diff." represents relative differences ((S3-BS)/BS, %).





**Figure 6**. Monthly-averaged daily maximum variation of shortwave direct aerosol radiative forcing at the top of <u>the</u> atmosphere from S3 and the relative changes due to water partitioning into OPM and non-ideality of the organic-water mixture during January and July of 2013. "<u>Rel.Abs. Diff.</u>" represents absolute differences (S3-BS); "Rel. Diff." represents relative differences ((S3-BS)/BS, %).





**Figure 7.** Correlation of water partitioning on SOA with (a) RH (b) temperature at Jinan in winter and Nanjing in summer, and the contribution from each SOA component to the total SOA increase. In plot (a) and (b), The sensitivity "Abs. Diff." represents the daily maximum change of SOA that is calculated as S3-BS. Color box represents RH in (a) and the hour in the day in (b) when daily maximum change of SOA occurred. In (c), the left axis represents contribution of each SOA component to the daily maximum SOA change due to water partitioning, and the right axis represents the concentration of each SOA component.



**Figure 8.** Sensitivity of SOA formation to temperature (TEMP), relative humidity (RH)water mixing ratio, and the temperature dependence parameter of SVP ( $\Delta H/R$ ) at Jinan (JN, first column) and Nanjing (NJ, second column).



**Figure 8**. Monthly-averaged impacts of water partitioning into OPM on SOA. "<u>Abs. Diff.</u>" represents absolute differences (S3-S2); "Rel. Diff." represents relative differences ((S3-S2)/S2, %).



**Figure 9**. Monthly-averaged impacts of non-ideality of the organics-water mixture on SOA. <u>"Abs.</u> The relative humidity is showing on the right side of y-axis. Diff." represents absolute differences (S3-S1); "Rel. Diff." represents relative differences ((S3-S1)/S1, %).