

*This paper uses a flow tube reactor to simulate the atmospheric oxidation potential of organic molecular markers emitted from burning peat and eucalyptus. The peat fires smoldered while the eucalyptus ones flamed. The semi-volatile organic markers were measured off-line using GC-MS or IC-PAD. The study was able to identify the reaction potential and the formation of semi-volatile organic markers that are common to chemical mass balance models and source apportionment studies.*

*The study topic is good for ACP. The results are timely because oxidation flow reactors (OFRs) are coming into vogue, and we need to better understand the results that OFRs produce and the potential air quality implications. With that, the study data have value and should be published. However, this particular study also has some serious limitations that need to be better addressed.*

AC: We thank the reviewer for the detailed and constructive review of our manuscript. Below, we first respond to general comments and then we address specific comments. All corrections were highlighted in the **marked** manuscript (Sengupta\_manuscript\_ACPD2019\_review\_marked\_Final).

### **General comments:**

*For one, the paper organization makes it seem like there are two separate studies, an emissions study and an OFR study.*

AC: We recognize the reviewer's concern and agree that the study does consist of two main parts (discussion of fresh and aged BB emissions). Due to the complexity of the data (EFs: fresh vs. aged, gas vs. particulate phase emissions from six different fuels) we prefer to preserve the current structure of the manuscript. Initially, we intended to present our results for fresh and aged emissions together. However, in order to increase coherency in narration we described the fresh emissions and corresponding changes after OFR oxidation separately.

*The emissions part of the paper is substantially less useful because all of the semivolatile markers were observed in past studies and reported about repeatedly. Additionally, the discussion on emissions lacks information about how these peat burning results are different than other forms of vegetation burning and what that may or may not mean for emissions reactivity. More critical thinking is warranted to help emphasize how the peat burning emissions are going to impact the SOA compared with other biomass fuels. Otherwise, it does not belong at the forefront of the paper currently. The best approach may be to emphasize the emissions work less and get to the OFR results sooner.*

AC: We agree with the reviewer that some of the semivolatile markers described in this study have been already presented in other studies (Mazzoleni et al., 2007; Yatavelli et al., 2017). On the other hand, we are convinced that our research provides new and very important data on fresh emissions from globally important biomass fuels, which have not been characterized extensively before. We added the importance of the emissions from analyzed fuels in the “Introduction” section (Pages 3-4, lines 71-79). The goal of this study was not to identify new BB organic compounds in fresh and aged emissions. Our research was mainly focused on analysis of organic compounds (with well-developed methods) and compare them between fresh and aged BB emissions from different fuels. In addition to polar organic compounds we also analyzed PAHs, alkanes, cycloalkanes, etc. Moreover, our next goal is to identify and quantify new BB compounds. These data will be presented in our forthcoming papers. We agree with the reviewer that it would be better if the OFR results appear sooner in the manuscript, however, as we mentioned in our previous comment, because of the inherent complexity of the data we are inclined to keep the current structure of our manuscript.

*The lack of replicate tests is another serious study limitation. It appears that standard deviations were taken from another set of tests and applied. This makes it difficult to judge the quality of the concentration data given here and calls into question how exactly the statistics were performed. It also complicates the interpretation of evidence showing what reacted or formed. This is a major limitation that needs to be addressed by providing the readers with sample population (N) and a detailed description of how the standard deviation was calculated and used to determine statistical significance and so on.*

AC: We recognize the lack of replicate analysis is a limitation of the current study. Unfortunately, due to resource (limited quantity of fuels available) and budget limitations of the project, we were able to perform only single-burning experiments with each fuel. However, in our previous BB experiments which were run with comparable conditions (Samburova et al., 2016; Yatavelli et al., 2017), we were able to perform triplicate burns for each fuel (including peats) and we argue that the uncertainties from our previous campaigns can be used for the current burning experiments. A detailed description was added to the section 2.5 The uncertainty of the emission factor values reported in the current manuscript includes: 1) burn to burn variability, 2) wall losses in the OFR, and 3) analytical uncertainty of the GC-MS method. In our previous campaigns, we clearly observed that the burn-to-burn variability is much higher than the wall losses and the analytical uncertainty of our method. There is no reason to believe that burn-to-burn variability, as well as other uncertainties in the current study are different from those observed in our previous studies. Therefore, we think that the best approach is to use the levels of uncertainty that we calculated for our previous combustion campaigns (calculated based on three replicate burns where a comparable experimental set up was used) and use these values for the present data set. For levoglucosan analysis with IC-PAD technique, we were able to calculate only analytical uncertainty.

*Finally, no new emissions compounds or SOA markers were measured that weren't part of the original compound suite. At the very least a set of chromatograms should be provided showing the raw emissions and the OFR effluent. That way we can see the major changes in these emissions.*

AC: We thank the reviewer for this proposition. We could not agree more that finding new SOA markers is really important in this area of atmospheric research. As mentioned in the current manuscript, we actually performed GC-MS full scan runs on our samples and we are currently working on the data for a future publication discussing compounds that fall outside of the current method.

### Specific comments:

1. *line 13: conserve use of the term 'significant' for describing statistics results.*

AC: We changed “significant” on “ prominent” (Page 1, Line 12). The word “significant” related to statistical results was also reviewed and changed through the text

2. *line 48: Biomass burning particles exhibit a variety of toxicological properties other than mutagenicity. It is worth being clear about that.*

AC: We agree, BB particle and gas exhibits wide toxicological properteies and mutagenicity is one of them. Hence, we have rephrased the sentence and added a few more refences on the toxicity of BB aerosols (Page 3, lines 47 – 49)

3. *lines 54-59: [MAJOR] Resolving the organic chemical composition of biomass burning particles is indeed challenging and improvements in speciation are needed. However, this particular study doesn't improve speciation technology in any way. It examines a common set of polar organic compounds and simulates their potential to photooxidize in the atmosphere. The suggestion is to focus on the reactivity of these molecules not on the improved speciation seeing that there are no methodological improvements in molecular level speciation being presented per se. After this study, we are still near 80% w/w for the unknown chemical fraction.*

AC: We appreciate reviewer's point and agree that this study is not novel in identification of new BB species. The goal of this research was to compare fresh and OFR aged BB emissions of known compounds, which have been already identified in previous studies. We deleted this part of the text from the main manuscript to avoid a wrong impression

that we also characterized unknown BB compounds (Page 3, lines 55-60). Also, we added a sentence (Pages 5-6, lines 132-135 of the marked manuscript) that clarifies the goal of this study.

4. *lines 61-72: This study uses peat as the main BB fuel. Be more specific about the importance of peat burning and why there needs to be a global emphasis on peat burning emissions*

AC: This is a good recommendation. Analysis of both fresh and aged peat burning emissions is indeed one of the strengths of this work. We added a paragraph that explains the importance of the peat emissions (Pages 3-4, lines 71-79).

5. *lines 80:81: I think the paper needs to be written around this sentence. This is what will make this study novel, important, and worth reading to many. Try and get to this point sooner. It would be better to couch this paper as one that attempts to examine the stability of commonly used polar atmospheric organic markers in an oxidizing environment.*

AC: We do agree that in our paper we attempted to investigate the stability of commonly identified polar organic compounds in laboratory oxidizing environment. However, we also wanted to analyze and discuss the difference in EFs of these compounds between different fuels as total (gas + particulate phases) and as well as individual gas and particulate phases.

6. *lines 81-93: Describe the advantages and disadvantages of each of these studies, how these pertain and give rise to the current study. Describe more than what these studies did, i.e, mention why these studies are important and are worthy of further discussion. Otherwise, there is not much value here.*

AC: This is a great suggestion. We added both advantages and disadvantages of the studies to the manuscript (Pages 4-5, lines 101-113).

7. *lines 95-96: All peats except Eucalyptus? Why not just focus on peats?*

AC: We understand the reviewer's recommendation to focus on peat fuels only. However, Eucalyptus fuel is ubiquitous and has a large contribution to recent Australian wildfire emissions. Therefore, presenting our results for eucalyptus emissions is timely and we prefer to keep them in our manuscript with respect to global emission perspectives. The importance of eucalyptus is now highlighted in section 2.1 (Page 7, lines 167 -171)

8. *lines 97-98: Describe why the focus here was isolated to polar compounds. It may have been more interesting to focus on the most reactive compounds in the aerosol mixture instead of presenting the study as an exercise in chemical accounting.*

AC: We understand the reviewer point. 84 polar compounds were selected and analyzed based on our established "polar compound" method, since there are no data how these compounds behave during OFR-aging in the complex mixture of real BB emissions.

9. *lines 135-140: At what depth were the tropical and Russian peats extracted?*

AC: The peat soils were extracted from the top 15cm of the soil profile, in accordance with convention and also to represent the depths at which soil combustion is most likely to occur. More details were added to the Experimental section (Page 7, lines 163-165)

10. *lines 176-179: Briefly mention the mass of fuel used per test and how the fuels were configured for burning.*

AC: Masses of all fuels were added to the Experimental section and the description of fuels' configuration during burning experiments was also added (Page 7, lines 179-180)

11. *lines 203-207: Figure 1 does not indicate which on-line instruments were used. These instruments should be described. Please explain how they were used. If they aren't being*

*used as part of this study please explain why. There is no modified combustion efficiency evidence showing the levels of smoldering or flaming combustion for each fuel type. Perhaps that can be included here.*

AC: The following on-line instruments were used during our combustion experiments: Scanning Mobility Particle Sizer (SMPS), Photo Acoustic Soot Spectrometer (PASS-3), CO and CO<sub>2</sub> analyzer. More detailed description of the instrumental set-up can be found in our previous publication (Bhattarai et al., 2018). PASS and SMPS results will be used in our following manuscript, which will focus on optical properties of BB emissions. During this campaign, the CO<sub>2</sub> analyzer malfunctioned and thus we were unable to use/present the continuous CO<sub>2</sub> data and hence the modified combustion efficiency.

12. *lines 229-231 and lines 251-258: It is mentioned that a <sup>13</sup>C radio-labeled levoglucosan sample was added as internal standard. Wouldn't the <sup>13</sup>C isotope elute at the same time as the unlabeled compound? How was that accounted for in this study. The radio isotopes are typically applied for GC-MS. How did things work with the IC-PAD method knowing that the PAD is a non-specific detector and that the radio isotope may interfere?*

AC: Thank you for catching this error. Levoglucosan-d7 was used as an internal standard for the quantitative analysis of levoglucosan with GC/MS. For the IC-PAD analysis of levoglucosan (see comment #17) the filters were not spiked with any internal standards. This error was corrected (**Page 8, lines 194-195**).

13. *lines 236-237: What is the value of reporting these separately considering that the equilibrium partitioning changes and was specific to the sampling conditions used for this particular study? Please discuss briefly. Also, was any thermodynamics-based check performed on how well the equilibrium partitioning was measured here? Such a relatively straightforward check on select compounds can verify that the sampling worked as planned.*

AC: This is definitely an interesting scientific question. We agree that the equilibrium partitioning can change under different sampling conditions. However, a compound's phase also affects its reaction rates. Thus, we believe it is important to report gas and particle concentrations as they were observed during the experiments. We did check whether the observed partitioning is thermodynamically reasonable – it is for most compounds, at least qualitatively. A quantitative assessment, however, is complicated for the heavy and light compounds due to the very low concentrations in one of the phases.

We are working on resolving these issues, but the complexity of this process is well beyond the scope of this manuscript .

14. lines 275-280: [MAJOR] Please clarify how the SDs were calculated. It's unclear where the standard deviations (SDs) came from and the reader should know exactly what and how many experiments were used to report SD values. The suggestion that the SDs came from another study and therefore can be applied here is questionable analytical-chemical practice. SDs should be experiment-specific, not just assumed and carried-over from study to study. At the very least, multiple injections of the same extract should be performed. What is meant by 'similar' fuels? This comment applies to all of the experimental data being reported.

AC: In the section 2.5 we added a description how the SDs were calculated for the present data set (Page 11, lines 280 – 293). In our analytical method we have replicates for few samples with multiple injections that we have used for calculating analytical uncertainties. However, burn-to-burn variability during the combustion experiments governs the overall uncertainties and hence we have included that variability calculated from the data collected in our previous combustion campaigns. It is worth mentioning that analytical uncertainties are also discussed along with burn-to-burn variability. The uncertainties for the peat combustion experiments were computed based on the results for Alaskan peat emissions (triplicate burns), while the uncertainties for eucalyptus were calculated using our previous Cheatgrass results (triplicate burns). The expression “Similar fuels” refers to similar combustion conditions (smoldering type peat combustion vs. flaming type eucalyptus combustion) (Page 13, line 320)

15. lines 282-328: This information about methoxy-phenol emissions from biomass burning is already available 'ad naseum' in the literature and can be further consolidated. Instead of focusing on what we already know, examine how the peat emissions are different (if they indeed are) and how these differences may be important to the atmospheric processing and oxidation. Knowing more about the extent of smoldering and flaming combustion can help properly develop this discussion.

AC: We have this discussion (differences in EFs between fuels) in our manuscript and we also compared our EFs with EFs from other studies. (Section 3.1.1).

16. lines 381-508: [MAJOR] Again, virtually all of this chemical information is already available. Much of this can be condensed into a brief paragraph or two unless there is something salient and different for the fuels being examined here that merits further attention. Suggestions: Develop one section that discusses 'Polar organic chemistry' of these emissions. Move the figures of individual compound emissions into supporting information. Develop a new emission figure that combines everything by compound class and shows the reader some sort of chemical mass balance so they can understand the aerosol fraction of interest (more like Figure 3, but include a mass balance). Only if needed, focus on some of the unique chemical or physical state attributes of individual compounds emitted during burning of these particular fuels. In other words, if the chemical structure or physical state of a compound is important to understanding its oxidation then discuss these attributes and the results. This requires a reorganization of the paper as some of these results are presented elsewhere.

AC: We really appreciate reviewer's recommendation of creating a new section and figures that exhibit a mass balance. We have addressed that issue in the general comment section too. Here, we give an example that explains why attaining closure in mass balance with our results is a difficult task and beyond the scope of this manuscript.

For example, aromatic acids can be produced during oxidation of methoxyphenols (reported in this study) and also from oxidation of PAHs and substituted PAHs (Wang et al., 2007) (not reported). Similarly, monocarboxylic acids can form by fragmentation (during oxidation) of compounds, form the same homologous series, and at the same time they also can be derived from oxidations of alkanes and alkenes. Because of the chemical complexity, it is very challenging to derive adequate mass balance results based only on analyzed polar compounds without considering other species.

17. line 509: It would be interesting to learn if any of the levoglucosan was found in the gas phase. Levoglucosan elutes from a GC column with and without derivatization. Was there any indication of levoglucosan in the gas-phase?

AC: For the quantitative analysis of levoglucosan we selected IC-PAD analysis over the GC-MS technique that includes BSTFA-derivatization of this polar compound. Due to the high levels of levoglucosan in the extracts, the MS detector was oversaturated; that made it impossible to perform adequate quantitative analysis of this compound. Therefore, further investigation of levoglucosan in the gas phase emissions was not feasible.



18. lines 524-530: *The important part of the study begins here. Figure 3 is a good figure that captures the essence of what was tested, measured, and and accomplished.*

AC: We thank the reviewer for this recognition

19. Figure 3: *Again, are these error bars taken from other studies? If so, please remove them from the figure. One idea is to combine all of the peat samples measured here and calculate a standard deviation and use that to report global error for experimental peat burning. That can be added as a panel to one of the figures or presented separately.*

AC: The uncertainties calculations are added to the text (**Page 11, lines 280-293**)

20. lines 605-614: *The importance of associating meaningful error with concentration data cannot be understated. For example, the error associated with levoglucosan measurement using GC-MS is approximately 20%. Although, the combustion test error can bring this value closer to 30%-50%. If we knew the exact error associated with repeating these peat experiments, we'd have a better understanding of this decrease being real or not. This may also explain why the same compound appears to be either formed or degraded for some fuels (see lines 631-632 and lines 691-699 (e.g., hexanoic acid)).*

AC: This is an interesting point. However, we believe that this is not only due to the error assignment. Hexanoic acid can be formed by oxidation of alkanes and decomposition of high molecular weight mono carboxylic acids (McNeill et al., 2008) This is why it is difficult to predict the fate of this compound by just looking at a matrix of polar compounds. Moreover, the occurrence of multiple generation of oxidations inside OFR (Li et al., 2015) leads to further complexity and as a result the same compounds can be either formed or degraded during oxidations of emissions from different fuels.

21. line 696: *Please describe the type of statistical tests being applied here and how they are being applied. Please report the sample population (N). Additionally, provide more information about how the error was calculated and applied to conduct the statistics tests.*

AC: We have described earlier how we calculated SD at the end of section 2.5 (**Page 11, lines 280-293**). In order to evaluate changes associated with oxidation experiments, we need to perform "Pairwise-T-test" with fresh and aged samples. However, in our case we

have computed the uncertainties from our previous campaign with only fresh samples, propagated that to both fresh and aged samples, and we refrained from performing pairwise T-test for this analysis. We agree with the reviewer's perspective that without replicate burns, we need to be conservative while describing some change as "statistically significant/insignificant". We have deleted such strong words to make the message less categorical. Changes in the main manuscript related to this comment are highlighted (**Lines 621, 681, 733, 747, 750**).

22. *lines 724-733: Please describe the criteria for a "top" contributing compound. Was its concentration?*

AC: The reviewer is correct, the "top" contribution compounds refer to compounds with highest levels of concentration (EFs). We rephrased the figure caption (**Page 35, lines 762-763**).

23. *lines 736-737: Isn't it also possible that the combustion was different?*

AC: We agree with the reviewer and added the "nature of combustion" to the text (**Page 36, line 776**).

## References:

- Bhatarai, C., Samburova, V., Sengupta, D., Iaukea-Lum, M., Watts, A.C., Moosmüller, H. and Khlystov, A.Y., 2018. Physical and chemical characterization of aerosol in fresh and aged emissions from open combustion of biomass fuels. *Aerosol Science and Technology*, 52(11), pp.1266-1282.
- Li, R., Palm, B. B., Ortega, A. M., Hlywiak, J., Hu, W., Peng, Z., Day, D. A., Knote, C., Brune, W. H., de Gouw, J. A. and Jimenez, J. L.: Modeling the Radical Chemistry in an Oxidation Flow Reactor: Radical Formation and Recycling, Sensitivities, and the OH Exposure Estimation Equation., *J. Phys. Chem. A*, 119(19), 4418–4432, doi:10.1021/jp509534k, 2015.

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- McNeill, V. F., Yatavelli, R. L. N., Thornton, J. A., Stipe, C. B. and Landgrebe, O.: Heterogeneous OH oxidation of palmitic acid in single component and internally mixed aerosol particles: Vaporization and the role of particle phase, *Atmos. Chem. Phys.*, 8(17), 5465–5476, doi:10.5194/acp-8-5465-2008, 2008.
- Samburova, V., Connolly, J., Gyawali, M., Yatavelli, R. L. N., Watts, A. C., Chakrabarty, R. K., Zielinska, B., Moosmüller, H. and Khlystov, A.: Polycyclic aromatic hydrocarbons in biomass-burning emissions and their contribution to light absorption and aerosol toxicity, *Sci. Total Environ.*, 568, 391–401, doi:10.1016/j.scitotenv.2016.06.026, 2016.
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- Yatavelli, R. L. N., Chen, L.-W. A., Knue, J., Samburova, V., Gyawali, M., Watts, A. C., Chakrabarty, R. K., Moosmüller, H., Hodzic, A., Wang, X., Zielinska, B., Chow, J. C. and Watson, J. G.: Emissions and Partitioning of Intermediate-Volatility and Semi-Volatile Polar Organic Compounds (I/SV-POCs) During Laboratory Combustion of Boreal and Sub-Tropical Peat, *Aerosol Sci. Eng.*, 1(1), 25–32, doi:10.1007/s41810-017-0001-5, 2017.

*Reviewer 2. This study quantified a number of compounds produced by burning six different fuels, most of them peat, and examined the effect of aging of the biomass burning smoke by reactions with OH (and to some extent ozone) in an oxidation flow reactor (OFR). The authors report emission factors (EFs) for all the compounds as well as the change in the effective EFs after OFR.*

AC: We thank the reviewer 2 for the constructive review of our paper. First we responded to major comments and then we addressed specific comments. All corrections were highlighted in the manuscript.

*Reviewer 2: Major comments 1. The reported dataset is comprehensive but it is not very clear to the reader what one can do with these data. While the decrease in the EFs can likely be modeled with known OH rate constants, the increasing EFs are hard to interpret without knowing whether additional products came from gas-ozone chemistry or from heterogeneous oxidation of particles by OH. There are some hints of this discussion in the paper, for example, on line 554. It would be good to expand this discussion.*

AC: We acknowledge the reviewer's concern regarding interpretation of the reported EFs for the aged samples. For example, for some fuels (Malaysian and Malaysian agricultural peat) hexanoic acid levels (EFs) were higher after the OFR-oxidation, while for Pskov peat there was an obvious decrease after the OFR (Figure 4c, Page 33). The formation of hexanoic acid after the OFR can be explained by the oxidation of alkanes in fresh BB emissions. However, to confirm this theory, more analyses with a simplified set of compounds are needed. Considering the fact that we analyzed only 84 specific polar compounds in a very complex BB mixture, it is difficult to interpret the formation of different products without speculation. We also acknowledge that the oxidation on the filters during the sampling can affect chemistry of the collected BB aerosols. However, considering high concentration of OH radicals and residence time of BB emissions in the OFR (Bhattarai et al, 2018), we believe that OH radical reactions in the gas-phase prevails over the gas-phase ozone chemistry and heterogeneous oxidation in the particulate-phase inside the OFR chamber. This discussion is presented on Page 36 and we also added one more sentence that highlights this issue (Page 10, line 246-247)

*Reviewer 2: Major comments 2. I always find it surprising how small the mass fraction of quantified compounds is despite clearly sophisticated analytical chemistry methods. Would it be possible to discuss what the rest of the compounds could be? Are there giant peaks in the GC data that are not reported because of the lack of standards?*

AC: We appreciate the reviewer's concern about quantifying only a small fraction of the total organic aerosol mass. In earlier studies, semivolatile organic compounds from biomass burning emissions were analyzed semi-quantitatively through both online (Fortenberry et al., 2018) and offline (Jen et al., 2019) sample collection techniques. Those studies also tried to establish mass balance closure of different compound classes but only with identified fraction of organic aerosol mass. However, no attempt was made to compare the unidentified compounds to the total organic mass. In our work, we performed full quantification of 84 compounds and we found those compounds account only for 5-7% of total organic aerosol mass.

Yes, the reviewer suggestion regarding other unknown species that can be present in the extracts (unidentified peaks on GCMS chromatograms) is absolutely right. We do observe some high-intensity peaks on full-scan GC-MS chromatograms of our samples. For example, NIST library search indicates that one of those peaks is ribose type sugars. Currently we are working on identification of these unknown species. Further results on this ongoing work will be presented in our forthcoming publication.

### Specific comments:

1. *Lines 126-133, 144-154: These sentences belong to either introduction or discussion. In the experimental section, it is better to focus on the specifics of the fuels used (such as water content, carbon content, etc.), not on the general importance of peatland ecosystems and eucalyptus forests*

AC: In the Introduction section we added the paragraph on the importance of peat fuels per reviewer suggestion (**Pages 3-4, lines 71-79**), however we think it is still better to keep some description of fuels in the Experimental section too.

2. *Line 165: it would probably help to know which molecule comes from which company.*

AC: The company names were added (**Page 8, lines 194-195**) and highlighted in the reviewed manuscript.

3. *Line 215: I do not think it is a sufficient argument. Reactions with OH are limited to a short time the mixture spends in OFR. In contrast, residual ozone comes out from OFR and has a chance to react with compounds on filter for the duration of sampling. Therefore, secondary oxidation of BB compounds on filter by ozone is quite possible and can only be ruled out if the residual ozone is low. The authors discuss these effects later in the paper on line 760.*

AC: Please see our response on the major comment 1.

4. *lines 61-72: This study uses peat as the main BB fuel. Be more specific about the importance of peat burning and why there needs to be a global emphasis on peat burning emissions.*

AC: We agree and we emphasized the importance of analyzed fuels in the reviewed manuscript (**Pages 3-4, lines 71-79**).

5. *Line 287: MW should be 124.14 g/mol*

AC: Typo with the error removed (**Page 14, line 328**)

6. *Line 360: missing right parenthesis*

AC: Parenthesis was added (**Page 17, line 399**)

7. *Line 492: replace )( with a semicolon*

AC: The problem with the reference was fixed (**Page 24, lines 529-530**).

8. *Lines 568, 660: you already used LMW on line 448, it would be better to define it there, or not use this abbreviation at all, as it is used only a few times.*

AC: Thank you for the suggestion. We corrected the text and use “low MW” and “high MW” everywhere in the text. The “MW” abbreviation is explained on **Page 1, line 29**.

9. *Line 664: error in  $2.4 \pm 0.001$  appears to be unrealistically small considering other errors reported in this paper.*

AC: The reviewer is right. This error can not be that small and was due to a typo. This typo was corrected (**Page 32, line 701**).

10. *Figure 4c: two panels have only one number on the X-axis, and ticks are too small making it hard to estimate the values from the figure (the tick comments applies to all similar figure).*

AC: Ticks were fixed for Figures 4a, 4b, 4c (**Pages 29, 31, and 33**).

11. *Figure 5: the text (on line 736) suggests that similar figures exists for other fuel types, however, they are not presented in the supporting information section*

AC: The figures (**Fig. S3 and Fig. S4**) were added to the SM.

## References:

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1 **Polar semi-volatile organic compounds in biomass burning emissions and their**  
2 **chemical transformations during aging in an oxidation flow reactor**

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9 **Abstract**

10 Semi-volatile organic compounds (SVOCs) emitted from open biomass-burning (BB) can  
11 contribute to chemical and physical properties of atmospheric aerosols and also may cause adverse  
12 health effects. The polar fraction of SVOCs is a prominent part of BB organic aerosols, and thus  
13 it is important to characterize the chemical composition and reactivity of this fraction. In this study,  
14 globally and regionally important representative fuels (Alaskan peat, Moscow peat, Pskov peat,  
15 Eucalyptus, Malaysian peat, and Malaysian agricultural peat) were burned under controlled  
16 conditions using the combustion chamber facility at the Desert Research Institute (DRI). Gas- and  
17 particulate-phase biomass-burning emissions were aged in an oxidation flow reactor (OFR) to  
18 mimic 5–7 days of atmospheric aging. Fresh and OFR-aged biomass-burning aerosols were  
19 collected on Teflon impregnated glass fiber filters (TIGF) in tandem with XAD resin media for  
20 organic compound (OC) speciation. The polar fraction extracted with dichloromethane and  
21 acetone was analyzed with gas chromatography mass spectrometry (GC-MS) for 84 polar organic  
22 compounds—including mono and dicarboxylic acids, methoxylated phenols, aromatic acids,  
23 anhydrosugars, resin acids, and sterols. For all these compounds, fuel-based emission factors (EFs)  
24 were calculated for fresh and OFR-aged samples. The carbon mass of the quantified polar  
25 compounds was found to constitute 5% to 7% of the total OC mass. High abundance of  
26 methoxyphenols (239 mg kg<sup>-1</sup> for Pskov peat; 22.6% of total GC-MS characterized mass) and resin  
27 acids (118 mg kg<sup>-1</sup> for Pskov peat; 14.5 % of total GC-MS characterized mass) was found in peat  
28 burning emissions (smoldering combustion). Concentration of some organic compounds (e.g.,  
29 tetracosanoic acid) with molecular weight (MW) above 350 g mol<sup>-1</sup> decreased after the OFR aging,  
30 while abundances of low MW compounds (e.g., hexanoic acid) increased. This indicated a  
31 significant extent of fragmentation reactions in the OFR. Methoxyphenols decreased after OFR  
32 aging, while a significant increase (3.7 to 8.6 times) in abundance of dicarboxylic acids emission



33 factors (EFs), especially maleic acid (10 to 60 times), was observed. EFs for fresh and ratios from  
34 fresh-to-aged BB samples reported in this study can be used to perform source apportionment and  
35 predict processes occurring during atmospheric transport.

36

37

38 **Keywords.** Biomass burning, organic aerosols, semi-volatile organic compounds (SVOCs), gas  
39 chromatography, mass spectrometry, polar organic compounds, oxidation flow reactor

40

## 41 1 Introduction

42 Biomass burning (BB), including both wildfires and prescribed burns, is a major source of  
43 carbonaceous aerosols in the atmosphere (Penner et al., 1991) and can contribute up to 75% of  
44 total atmospheric aerosol mass loading (Andreae et al., 2001; Park et al., 2007). These  
45 carbonaceous aerosols have significant impact on both regional and global radiative forcing  
46 (Ramanathan and Carmichael, 2008). BB emissions also can cause adverse health effects (Arbex  
47 et al., 2007; Regalado et al., 2006) because of the ~~mutagenic property~~ toxicological properties of  
48 particle-bound organic compounds (Chen et al., 2017; Pardo et al., 2020; Pavagadhi et al., 2013;  
49 Sigsgaard et al., 2015; Yang et al., 2010) ~~(for example, mutagenicity (Yang et al., 2010))~~.  
50 Therefore, comprehensive, molecular-level characterization of BB emissions is essential for  
51 understanding health effects. Such molecular characterization of BB carbonaceous aerosols in the  
52 atmosphere, however, is challenging as these aerosols are composed of tens of thousands of  
53 compounds (Goldstein and Galbally, 2007).

54  
55 ~~Current atmospheric chemistry models use a limited number of organic species because of the~~  
56 ~~complexity of atmospheric aerosol chemical composition and the lack of aerosol chemical~~  
57 ~~speciation data. Approximately 80% of BB organic mass emissions, especially aged emissions, are~~  
58 ~~not identified in such models (Bertrand et al., 2018; Jen et al., 2019), limiting the capabilities of~~  
59 ~~atmospheric organic aerosol modeling. Thus, improvement is needed in molecular level speciation~~  
60 ~~of both fresh and aged BB emissions for more accurate model estimations.~~

61  
62 Simulation of natural fires in a laboratory environment using a BB chamber is one way to  
63 characterize the chemical composition of BB emissions (Yokelson et al., 2003). A number of  
64 studies characterizing the molecular composition of combustion emissions from fuels that  
65 represent different geographical regions have been completed: temperate conifers (Oros and  
66 Simoneit, 2001a), deciduous trees (Oros and Simoneit, 2001b), grasses (Oros et al., 2006), and  
67 peats (Samburova et al., 2016; Iinuma et al., 2007). Akagi et al. (2011) compiled fuel-based  
68 emission factors (EFs) from different fuels from throughout the world, including the peatlands of  
69 south Asia, and found that burning condition (flaming/smoldering) can influence the EFs of  
70 individual compounds. These data have been used for modeling work in predicting ozone-forming  
71 potential and other air quality impacts (Alvarado et al., 2015). ~~Very few studies (e.g. (Samburova~~

72 et al., 2016) have focused on peat emissions. However, the importance of investigating the  
73 combustion products from burning peat soils is multifaceted. Peat soils, comprised predominantly  
74 of partially decomposed organic material, represent one-fourth to one-third of global terrestrial  
75 carbon and are under threat of increased fire activity in both boreal and tropical latitudes, areas of  
76 widespread peatland occurrence (Turetsky et al., 2015). In addition to the implications of peat fires  
77 for the global C cycle, local impacts from burning of peatlands include public health and safety  
78 problems from degraded to air quality, as well as ecological changes due to altered surface  
79 hydrology in low-relief areas (Watts et al., 2015). Most source apportionment studies, however,  
80 focused on characterization of fresh emissions and emissions of either particle-phase or gas-phase  
81 compounds.

82  
83 Significant changes in organic aerosol composition during atmospheric transport have been  
84 reported (Liu et al., 2017; Decker et al., 2019). These changes can impact local and regional air  
85 quality. Also, the role of Siberian peat burning in haze formation in the Korean peninsula (Jung et  
86 al., 2016) demonstrates the global impact of BB emissions and their atmospheric transport on  
87 regional air quality. Some laboratory studies found an increase in organic aerosol (OA) mass after  
88 photochemical aging (Ortega et al., 2013; Grieshop et al., 2009) while others observed a modest  
89 decrease (Bhattacharai et al., 2018). There is still limited data on evolution of chemical composition  
90 of primary organic aerosols (POAs) during atmospheric aging. Some laboratory experiments  
91 demonstrated degradation of levoglucosan (Hennigan et al., 2010; Kessler et al., 2010) and  
92 oxidation of methoxyphenols in the gas phase (Yee et al., 2013) and aqueous phase (Net et al.,  
93 2011). These studies have more mechanistic implications than quantifying gross change after  
94 atmospheric oxidation. The reactivity of a pool of organic compounds in a complex mixture such  
95 as BB emissions is expected to be different mechanistically from individual compounds. This  
96 necessitates the need for studies of the evolution of organic compounds through bulk molecular  
97 level characterization of BB emissions. Recently, Fortenberry et al. (2018) characterized the  
98 chemical fingerprints of aged biomass-burning aerosols (leaf and hardwood of white oak) by  
99 performing oxidation in a potential aerosol mass oxidation flow reactor (PAM-OFR) and chemical  
100 analysis with a thermal desorption aerosol gas chromatograph aerosol mass spectrometer (TAG-  
101 AMS). In this study, denuders were used to remove gases and particles were introduced to OFR to  
102 evaluate only the changes in particulate BB emissions during OFR oxidation. However, the

103 presence of both gas and particulate phase emissions in real BB emission and the partitioning of  
104 organic compounds in such a complex mixture can affect the reactivity inside the OFR and hence  
105 the fate of organic compounds during OFR aging. Bertrand et al. (2018) analyzed 71 organic  
106 compounds in BB emissions, sampled from a smog chamber, with high resolution time of flight  
107 mass spectrometry (HR-ToF-AMS). This study confirms that nitro-aromatic compounds are  
108 formed after OFR oxidation and that they can be used as SOA tracer. However, this study was  
109 focused on controlled wood (pellet) burning that is substantially different from wildland BB  
110 emissions. There is still a lack of understanding, however, regarding (1) major organic compounds  
111 emitted from BB (especially from peat fuels), (2) their roles in atmospheric photochemical  
112 reactions, and (3) what compounds are responsible for light absorption of fresh and aged BB  
113 emissions.

114  
115 In this study, emissions from laboratory combustion of six globally important fuels (Alaskan peat,  
116 Moscow peat, Pskov peat, Eucalyptus, Malaysian peat, and Malaysian agricultural peat) were  
117 quantitatively analyzed for more than 250 individual organic species, and analyses of 84 polar  
118 organic species is presented in this paper. BB emissions generated in a combustion chamber were  
119 run through the OFR, mimicking approximately 5 to 7 days of atmospheric oxidation (Bhattarai et  
120 al., 2018), and the OFR output was analyzed to characterize aged BB emissions. BB emissions  
121 were collected on filter and XAD media to identify distribution of organic species between the gas  
122 and particle phases. For the polar fraction of collected organic compounds, we quantitatively  
123 analyzed a total of 84 compounds (methoxyphenol derivatives, dicarboxylic acids,  
124 monocarboxylic acids, aromatic acids, resin acids, and anhydrosugars). In the analyzed  
125 anhydrosugars, we paid special attention to levoglucosan, a derivative from cellulose (Simoneit et  
126 al., 1999), since levoglucosan has been widely used as a molecular tracer of BB emissions  
127 (Bonvalot et al., 2016; Maenhaut et al., 2016). Methoxyphenols also have been used in source  
128 apportionment studies (Schauer et al., 2001; Schmidl et al., 2008b, 2008a). These source  
129 apportionment studies, however, haven't combined such a wide range of different groups in a  
130 single investigation. Here we provide a detailed targeted chemical analysis of both gas- and  
131 particle-phase BB emissions from the combustion of individual biomass fuels from diverse  
132 geographical locations for both fresh and aged emissions. The goal of this research was to  
133 characterize EFs of gas- and particulate-phase individual polar organic species are presented for

134 six compound groups of compounds (methoxyphenols, dicarboxylic acids, monocarboxylic acids,  
135 aromatic acids, anhydrosugars, and resin acids) and to analyze their fate after the-OFR oxidation.

136 In separate sections we discussed EFs obtained for fresh and OFR-aged BB samples. The fresh-  
137 to-aged ratio and top contributing organic species also are discussed. The comparison between  
138 fresh and OFR-aged BB emissions helps to understand the chemical evolution of BB plumes in  
139 the atmosphere and the obtained data can be used in future source apportionment and atmospheric  
140 modeling studies.

141

## 142 **2. Experiments**

### 143 **2.1 Fuel Description**

144 We selected six globally and regionally important BB fuels: Alaskan peat, Moscow peat, Pskov  
145 peat, Eucalyptus, Malaysian peat, and Malaysian agricultural peat. Five of these were peat fuels  
146 selected from different geographical locations, representing smoldering combustion and one  
147 (Eucalyptus) representing flaming combustion.

148 Peatland ecosystems, generally wetland or mesic ecosystems underlain by soils composed  
149 primarily of partially-decomposed biomass, contain mostly organic carbon and more than 20%  
150 mineral content, represent a vast terrestrial carbon pool, and are potentially vast sources of carbon  
151 flux to the atmosphere during wildfires that consume peat (Harden et al., 2000). Peatlands in high-  
152 latitude temperate and boreal regions are particularly vulnerable to increased fire-related carbon  
153 emissions resulting from climatic warming and increases in fire season length, while peatlands in  
154 low-latitude and tropical regions are threatened by factors such as deforestation for agriculture,  
155 urbanization, and drainage (Turetsky et al., 2015). We collected Alaskan peat samples from the  
156 upper 10 cm of soils within black spruce (*Picea mariana*) near crown forest (Chakrabarty et al.,  
157 2016). High-latitude and Eurasia samples here are from *Sphagnum*- and cotton grass- (*Eriophorum*  
158 spp.) dominated communities, collected from the Moscow (Odintsovo and Shatura districts) and  
159 Pskov regions of Russia. These regions are representative of oligotrophic peat bogs found widely  
160 across Siberia as well. Tropical peat in this study includes samples from two areas in Malaysian  
161 Borneo. One set of samples is from a *Dipterocarp*-dominated lowland forest with largely intact  
162 native land cover, while the second set is from a cleared agricultural area in the Kota Snamarahan

163 region. All peat soils were extracted from the top 15 cm of the soil profile. The mass of the dry  
164 peat fuels varied from 70-100 g. The peat fuels were burned inside a metal pot which contains  
165 thermally insulating materials at its bottom and sides.

166  
167 We selected Eucalyptus as a test fuel because of its prevalence across Australia and its important  
168 contribution to Australian wildland fires. In addition, economic losses and risk to life and property  
169 from fires in eucalypt forests are magnified by their proximity to both fire-prone ecosystems and  
170 large urban areas; often eucalypt-dominated stands form boundaries between these two land-use  
171 types. There are nearly 900 species of the genera *Eucalyptus*, *Corymba*, and *Angophora*, which  
172 collectively comprise woody plants known as eucalypts. Native to Australia, eucalypt-dominated  
173 forests cover nearly 92 million ha (Hills, W.E.; Brown, 1978). In addition, the fast and hardy  
174 growth characteristics of eucalypts have made them popular in warm ecoregions of Europe as well  
175 as North and South America, where they readily escape cultivation and become established,  
176 dominant community types near urban areas where they were originally introduced. Because of  
177 their high oil content, rapid and dense growth, and vegetative structure, eucalypts are highly  
178 flammable and contribute to high fire risk in areas where they occur (Goodrick and Stanturf, 2012).  
179 The mass of the Eucalyptus fuel burned (i.e., ~650 g) was higher than that of peat fuels and this  
180 fuel was arranged similar to a common arrangement for campfires.

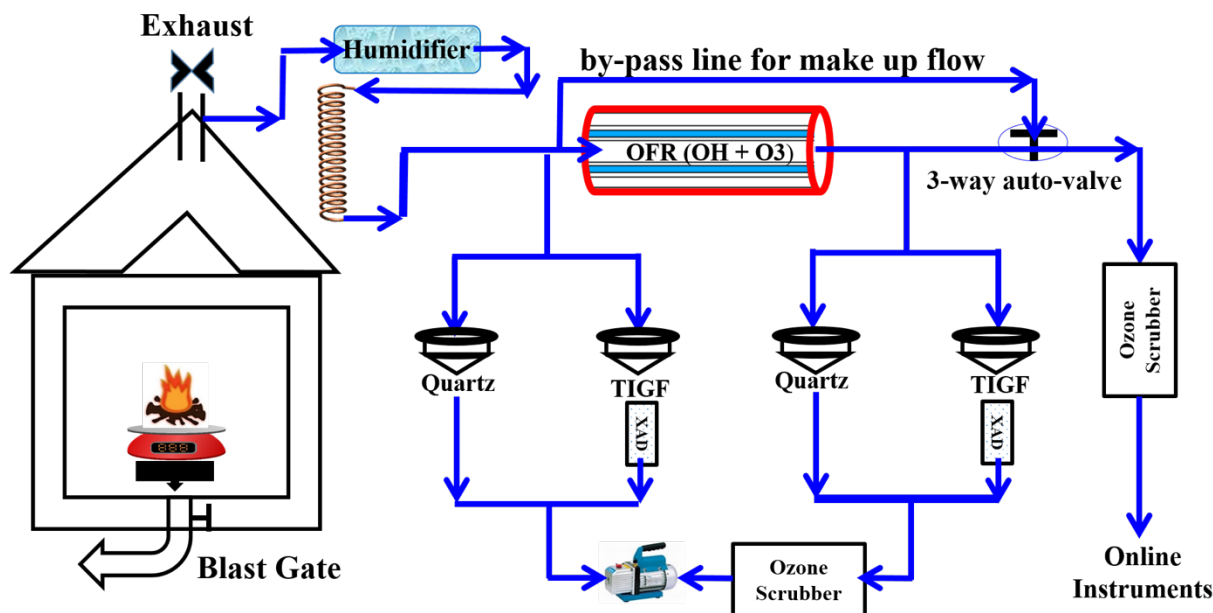
181

## 182 **2.2 Reagents and Materials**

183 We obtained high-performance liquid chromatography (HPLC) grade methanol and hexane from  
184 Fisher Scientific (Fair Lawn, NJ, USA) and used the following filters for sampling and further  
185 chemical analyzes: pre-fired (900 °C for 4 h) 47 mm diameter quartz-fiber filter (2500 Pallflex  
186 QAT-UP, Pall Life Sciences, Ann Arbor, MI, USA) for thermo-optical Elemental Carbon/Organic  
187 carbon (EC/OC) analysis, Teflon® filters (2500 Pallflex QAT-UP, Pall Life Science, Ann Arbor,  
188 MI, USA) for gravimetric particulate matter (PM) mass analysis, and Teflon-impregnated glass  
189 fiber (TIGF) 47 mm diameter filters (Fiber FilmT60A20, Pall Life Sciences, Ann Abor, MI, USA)  
190 for organic analysis. We purchased the following deuterated internal standards from Cambridge  
191 isotope laboratories (Tewksbury, MA, USA) and CDN isotopes (Pointe-Claire, Quebec, Canada):  
192 hexanoic-d11 acid, , succinic-d4 acid, decanoic-d19 acid, adipic-d10 acid, suberic-d12 acid,

193 homovanillic-2,2-d<sub>2</sub> acid, myristic-d<sub>27</sub> acid, heptadecanoic<sub>33</sub> acid, oleic-9,10-d<sub>2</sub> acid  
194 tetradecanedioic-d<sub>24</sub> acid (CDN Isotopes, Quebec, Canada) and benzoic-d<sub>5</sub> acid, levoglucosan-  
195 d<sub>7</sub> and cholesterol-2,2,3,4,4,6-d<sub>6</sub> (Cambridge Isotope Laboratories, Inc., MA, USA)

196  
197  
198 **2.3 Biomass Burning (BB) Experiments**



199  
200 **Figure 1.** Desert Research Institute (DRI) biomass burning (BB) facility with oxidative flow  
201 reactor (OFR) and flow setup.

202  
203 BB experiments were conducted using DRI's BB facility for combustion of the selected fuels  
204 under controlled conditions. A close replicate of this facility was described previously (Tian et al.,  
205 2015), and a detailed description of the experimental setup was presented elsewhere (Bhattarai et  
206 al., 2018; Sengupta et al., 2018).

207  
208 We mixed laboratory-generated BB emissions with humidified zero air (Airgas Inc., Sparks, NV,  
209 USA) using 4 m long spiral copper tubing (12.7 mm OD). Before it was mixed with the BB  
210 emissions, the zero air was humidified by bubbling through Nano-pure water in a glass 500 mL  
211 volume impinger. The flow rate was controlled with a mass flow controller (810C-CE-RFQ-1821,  
212 Sierra Instruments, Monterey, CA, USA). An oxidation flow reactor (OFR) (Aerodyne Research

213 Inc., Billerica, MA, USA) was used to mimic approximately seven days of equivalent atmospheric  
214 aging (Bhattacharai et al., 2018). The OFR consisted of an alodine-coated aluminum cylinder (46 cm  
215 length and 22 cm diameter) with an internal volume of 13.3 L. Two sets of lamps emitted UV  
216 radiation at wavelengths of 185 and 254 nm (Atlantic Ultraviolet Corporation, Hauppauge, NY,  
217 USA) in the OFR to produce ozone and OH radicals (Li et al., 2015). UV irradiance in the OFR  
218 was quantified using a photodiode detector with a wavelength range of 225 to 287 nm  
219 (TOCON\_C6; Sglux GmbH, Berlin, Germany). Ultra-high-purity nitrogen (Airgas Inc., Reno,  
220 NV, USA) was used to purge the UV lamp compartments to prevent the lamps from overheating.  
221 A probe that monitored relative humidity and temperature inside the OFR (from Aerodyne Inc.,  
222 MA, USA) was mounted toward the outlet side of the OFR. A detailed characterization of the  
223 OFR—such as particle loss, OH production rate, and time scales of various processes—can be  
224 found in Bhattacharai et al. (2018).

225  
226 The duration of smoldering combustion experiments ranged from 69 to 255 min, whereas the  
227 average duration of flaming combustion experiments was 50 min. During all experiments, both  
228 fresh (directly from the chamber) and aged (oxidized in the OFR) emissions were continuously  
229 collected on a TIGF filter (for particle phase) followed by an XAD cartridge (for gas phase) for  
230 detailed chemical speciation. We used several online instruments to characterize gas- and particle-  
231 phase pollutants (see Fig. 1). Simultaneous collection of samples for thermal optical carbon  
232 analysis on quartz fiber filters (Pall-Gelman, 47 mm diameter, pre-heated) was conducted, but only  
233 for Eucalyptus and Malaysian peat. The online instruments alternated every 10 min between  
234 sampling fresh and aged emissions using a computer-controlled valve system. A **description of all**  
235 **instruments and the detailed experimental set up can be found elsewhere (Bhattacharai et al., 2018).**

236  
237 We employed a bypass flow to keep the flow from the BB chamber and through the OFR constant  
238 when online instruments switched between sampling fresh and aged emissions. To protect online  
239 instruments from high ozone concentrations produced in the OFR, ozone scrubbers were installed  
240 in front of the instruments' inlets. The ozone scrubbers were loaded with charcoal followed by  
241 Carulite 200 catalyst (Carus Corp., Peru, IL, USA). There were no ozone scrubbers before the  
242 filter-XAD set up, which could cause further oxidation of organic compounds on filter surfaces  
243 during sampling. The reaction rates between organics and ozone, however, are orders of magnitude



244 lower than OH oxidation reactions (Finlayson-Pitts and Pitts Jr, 1999). Therefore, we assumed that  
245 reactions with OH radicals were primarily responsible for changes in organic compounds  
246 associated with fresh gas and particulate emissions. We also acknowledge that the ozone-oxidation  
247 on the filters during the sampling can affect chemistry of the collected BB aerosols (section 3.4).

248

#### 249 **2.4. Organic and Elemental Carbon (OC/EC) Analysis**

250 Emissions from the combustion of two fuels (Eucalyptus and Malaysian peat) were sampled with  
251 quartz-fiber filters, collected simultaneously with TIGF filters, for both fresh and aged BB aerosols  
252 (Supplementary Material, Fig. S1). Punches (area = 1.5 cm<sup>2</sup>) from these quartz filters were  
253 analyzed with a thermal-optical carbon analyzer (Atmoslytic Inc., Calabasas, CA, USA) following  
254 the IMPROVE protocol (Chow et al., 1993, 2004) for total organic carbon (OC<sub>Total</sub>) and elemental  
255 carbon (EC) mass.

256

#### 257 **2.5 Analytical Methodology for GC-MS**

258 We extracted filter and XAD samples for GC/MS analysis (SI Table S1) yielding concentrations  
259 of 84 polar organic compounds. In addition, levoglucosan concentrations were determined using  
260 ion chromatography coupled with a pulsed amperometric detector (IC-PAD). Prior to the  
261 extraction, sampled TIGF filters and XAD-resin cartridges were spiked with deuterated internal  
262 standards (see “Reagents and Materials” section). The TIGF filters and XAD cartridges were  
263 extracted separately with an accelerated solvent extractor (Dionex ASE-300, Sunnyvale, CA,  
264 USA) at the following conditions: 80° C temperature, 250 mL extraction volume, and subsequent  
265 extraction with dichloromethane and acetone. The XAD and filters were treated separately to  
266 evaluate the speciation of gas- and particle-phase semi-volatile polar compounds. The extracts  
267 were concentrated with a rotary evaporator (Buchi-R124, Switzerland), filtered using 0.2 µm pore  
268 size syringe filters (Thermo Scientific, Redwood, TN, USA), and pre-concentrated with nitrogen  
269 to a volume of 4 mL. Then we split the extracts into two fractions. One fraction was transferred to  
270 2.0 mL volume deactivated glass maximum recovery vials (Waters Corporation, Milford, MA,  
271 USA), pre-concentrated to 50 µL volume under ultra-high-purity nitrogen (Airgas, Reno, NV,  
272 USA), and derivatized with N,O-bis-(trimethylsilyl) trifluoroacetamide (BSTFA with 1% of  
273 trimethylchlorosilane; Thermo-Scientific, Bellefonte, PA, USA) and pyridine as described

274 elsewhere (Rinehart et al., 2006). Derivatized samples were analyzed by electron impact ionization  
275 using a Varian CP-3400 gas chromatograph with a CP-8400 auto-sampler and interfaced to a  
276 Varian 4000 ion trap mass spectrometer (Varian Inc. Palo Alto, CA, USA). The second fraction of  
277 non-derivatized extracts was kept for further analysis of non-polar organic species (e.g., alkanes  
278 and PAHs), and those results will be presented in future publications.

279  
280 Since, we have not performed any replicate burning experiments in this study, we used uncertainty  
281 values from our previous work (Yatavelli et al., 2017) as follows. We took mean ( $\mu_c$ ) and standard  
282 deviation ( $\sigma_c$ ) values for all 84 compounds from Alaskan peat (3 replicates) and Cheatgrass (3  
283 replicates) combustion and calculated fractional uncertainties ( $f_c$ ) using equation 1.

$$284 \quad f_c = \frac{\mu_c}{\sigma_c} \quad \text{Eq. 1}$$

285 These fractional uncertainty values are multiplied by the emission factors for each compound (both  
286 fresh and aged) from the current study to obtain individual uncertainties or standard deviation (SD)  
287 values as demonstrated in equation 2.

$$288 \quad SD_c = f_c \times EF_c \quad \text{Eq. 2}$$

289 The group uncertainties for each different subclass were computed from individual uncertainties  
290 by applying standard propagation of error method (root sum of squares).

291 Considering the different nature of combustion, the fractional uncertainties derived from Alaskan  
292 peat combustion were applied to emissions from all peat fuels and the uncertainties obtained from  
293 cheatgrass combustion were applied to emissions from Eucalyptus combustion.

294

## 295 **2.6. Levoglucosan Analysis**

296 Portions of quartz filters collected for OC/EC analysis also were used for quantitative analysis of  
297 levoglucosan concentration with IC-PAD. Prior to the analysis, quartz filters were extracted with  
298 15 ml of deionized water (18.2 M $\Omega$ ), sonicated for one hour, and refrigerated overnight. The  
299 column temperature for IC was 25 $^\circ$  C. Analytes along with a mixture of two eluents (48%  
300 hydroxide solution and 52% deionized water) were passed through the IC column with a 0.4 ml  
301 min $^{-1}$  eluent flow rate and detected using an electrochemical detector. See Chow and Watson  
302 (2017) for details. Uncertainties or standard deviations associated with levoglucosan results  
303 indicate only analytical uncertainties of the IC-PAD method.

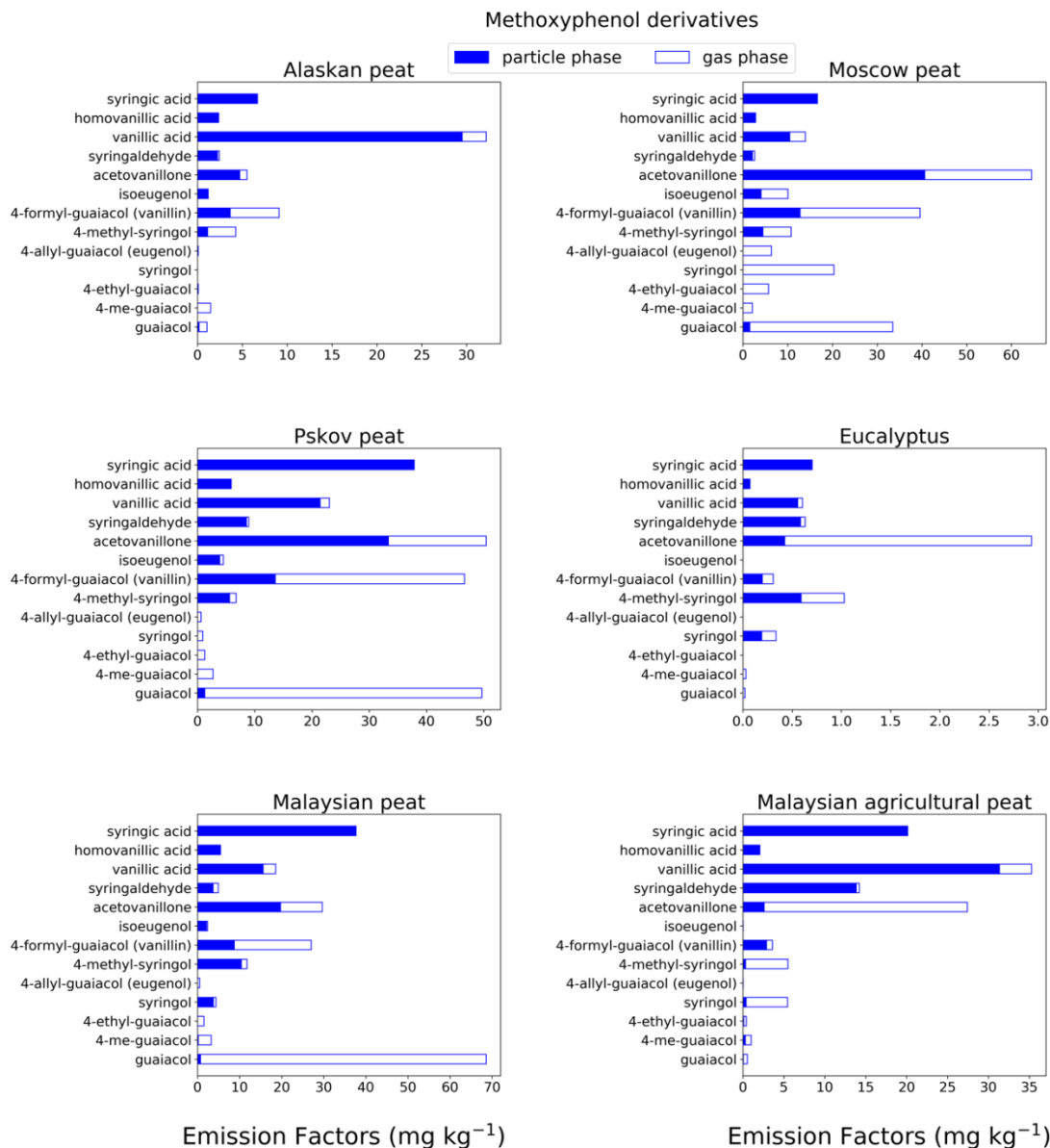
### 304 **3. Results and Discussion**

305

#### 306 **3.1. Gas- and Particulate-Phase Emission Factors**

307 Organic compounds (84 in total) in fresh emissions identified and quantified in this study were  
308 assigned to six major groups (Table S1): methoxyphenol derivatives, dicarboxylic acids, mono-  
309 carboxylic acids, aromatic acids, resin acids, and levoglucosan. First, we report individual  
310 emission factors (EF) belonging to a particular group calculated by summation of gas- and particle-  
311 phase EFs of individual compounds. Relative abundance of these compounds **is** reported next  
312 followed by a comparison of the contributions of each group ( $EF_{\text{group}}$ ) among fuels and a  
313 comparison with previously reported results.

314



316

317 Figure 2a. EFs for methoxyphenols in both particulate phase (solid bars, filter samples) and gas-  
 318 phase (open bars, XAD samples) from fresh biomass burning emissions for six different fuel types.  
 319 We did not burn fuels in replicates, and standard deviations (SD) were calculated based on replicate  
 320 analysis of emissions from **similar** fuels (with identical experimental conditions) during our  
 321 previous combustion campaigns (Yatavelli et al., 2017) where SD ranged between 9.7 and 22%  
 322 for methoxyphenol derivatives.

323 Methoxyphenols are key compounds in BB smoke since they constitute from 20 to 40% of total  
324 identified organic aerosol mass (Hawthorne et al., 1989; Yee et al., 2013). For this reason, these  
325 compounds are considered potential markers for wood combustion (Schauer et al., 2001) and have  
326 been used as probable biomarkers to determine human exposure to BB emissions (Simpson and  
327 Naeher, 2010; Dills et al., 2006). Our analysis of 13 methoxyphenols (Fig. 2a, Table S1) showed  
328 that guaiacol (MW = 124±12 g mol<sup>-1</sup>) was the major contributor to EFs of the measured  
329 methoxyphenols in Moscow peat (33.5±3.3 mg kg<sup>-1</sup>), Pskov peat (49.7±4.8 mg kg<sup>-1</sup>), and  
330 Malaysian peat (68.6±6.7 mg kg<sup>-1</sup>). Syringol, another methoxyphenol commonly found in BB  
331 emissions (Schauer et al., 2001), had the highest EF for Moscow Peat fresh emissions (20.4±2.7  
332 mg kg<sup>-1</sup>), while for the other fuels, the EF was much lower (0–5.5 mg kg<sup>-1</sup>). EFs for syringic acids  
333 (MW = 198 g mol<sup>-1</sup>) were in the range of 0.06–37.9 mg kg<sup>-1</sup> for all fresh emissions. Syringols are  
334 generally not formed during pyrolysis of coniferous lignin, but during pyrolysis of deciduous  
335 lignin, where both guaiacols and syringols are formed (Mazzoleni et al., 2007). Presence of both  
336 guaiacol and syringol moieties in fresh emissions indicates that the part of the plant material that  
337 was responsible for peat formation was probably from deciduous trees, and this signature of  
338 deciduous trees from peat burning emission is irrespective of geographical origin of those peats  
339 (also shown by Schauer et al., 2001). Acetovanillone, vanillin, and vanillic acid also were observed  
340 in fresh emissions with high abundance (5–50 mg kg<sup>-1</sup>). For example, vanillin is an abundant  
341 methoxyphenol in the fresh emissions from Pskov peat (46.7±5.4 mg kg<sup>-1</sup>) which contributed 4.4%  
342 of the total mass of the 84 analyzed compounds.

343  
344 **Low MW** methoxyphenols (e.g., guaiacol) are expected to be found in the gas phase (Yatavelli et  
345 al., 2017), in close agreement with our results. For example, guaiacol and substituted guaiacols  
346 were mostly present in gas phase (82–100%) for emissions from the combustion of different fuels  
347 (Fig. 2a, Table S1). With the addition of more oxygenated functional groups to a molecule, and  
348 thus with **MW** increase, the equilibrium gas-particle partitioning of the compound tends to shift  
349 toward the particulate phase, which also was confirmed by our results (e.g., for acetovanillone, a  
350 keto form of lignin derivative, from Malaysian peat combustion, 33.5% of its mass was found in  
351 gas phase; for more oxygenated syringic acid, 99% of its mass was found in particulate phase  
352 emissions from the same fuel).

353 The highest methoxyphenol  $EF_{\text{group}}$  from combustion of all fuels was observed in the fresh Pskov  
354 peat (Fig. 3a) emissions ( $239 \pm 11 \text{ mg kg}^{-1}$ ). For Moscow peat, which was sampled close to the  
355 geographical region of Pskov peat, the  $EF_{\text{group}}$  of methoxyphenols was  $229 \pm 10 \text{ mg kg}^{-1}$  (Fig. 3a),  
356 very similar to that for Pskov peat. The methoxyphenol  $EF_{\text{group}}$  for peat samples were in the range  
357 of 66 to  $239 \text{ mg kg}^{-1}$  (Fig. 3a) for our 13 analyzed compounds. A previous study analyzed for 30  
358 different compounds (Schauer et al., 2001) and consequently found a larger value  $EF_{\text{group}}$  of up to  
359  $1330 \text{ mg kg}^{-1}$ , at least partially a result of the larger number of compounds analyzed. Formation  
360 of methoxyphenols during biomass combustion is mainly because of pyrolysis of lignin (e.g.,  
361 Simonelt et al., 1993). Lignin, an essential biopolymer of wood tissue, is primarily derived from  
362 three aromatic alcohols: p-coumaryl, coniferyl, and sinapyl alcohols (Hedges and Ertel, 1982).  
363 Lignins of hardwoods (angiosperms) are enriched with products from sinapyl alcohol; softwoods  
364 (gymnosperms) instead have a high proportion of products from coniferyl alcohol with a minor  
365 contribution from sinapyl alcohol; grasses have mainly products from p-coumaryl alcohol. The  
366 relative proportions of these bio-monomers vary considerably among the major plant classes  
367 (Sarkanen and Ludwig, 1971), reflected in our total emission factors estimate for 13  
368 methoxyphenols.

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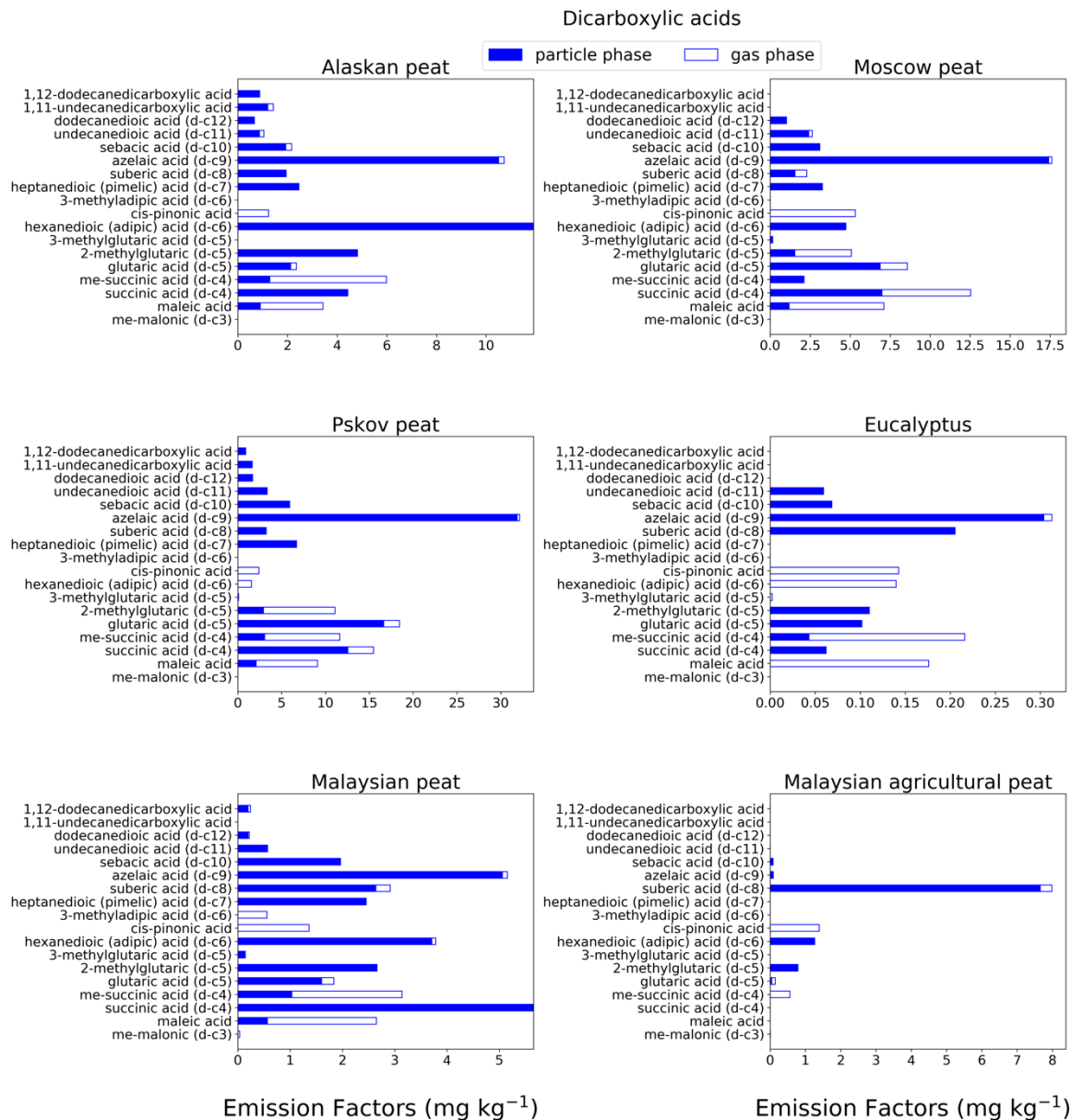
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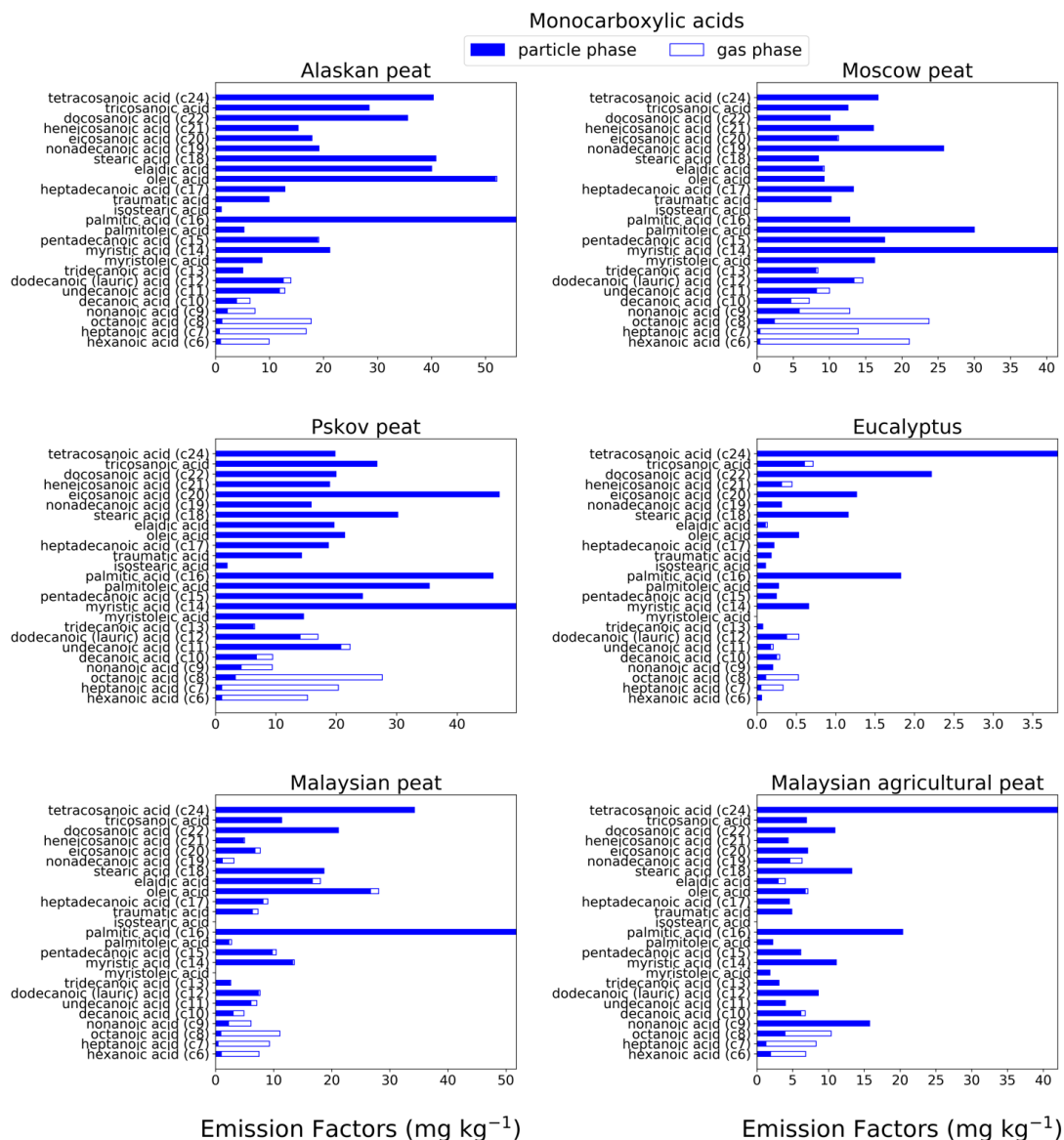


380

381 Figure 2b. EFs for dicarboxylic acids in both particulate phase (solid bars, filter samples) and gas-  
 382 phase (open bars, XAD samples) from fresh biomass-burning emissions for six different fuel types.  
 383 We did not burn fuels in replicates, and standard deviations (SD) were calculated based on replicate  
 384 analysis of emissions from similar fuels (with identical experimental conditions) during our  
 385 previous combustion campaigns (Yatavelli et al., 2017) where SD ranged between 10 and 17% for  
 386 dicarboxylic acids.

387 Dicarboxylic acids play a significant role in the atmospheric organic aerosols budget (Samburova  
388 et al., 2013; Yatavelli et al., 2017) via secondary organic aerosol formation that either changes  
389 radiative forcing directly, or indirectly by acting as cloud condensation nuclei (Kawamura and  
390 Bikkina, 2016). The  $EF_{\text{group}}$  for dicarboxylic acids (Fig. 3b) varied among the fuels with the highest  
391  $EF$  for fresh Pskov peat samples ( $123 \pm 10 \text{ mg kg}^{-1}$ ) and with the lowest for Eucalyptus ( $1.5 \pm 0.1$   
392  $\text{mg kg}^{-1}$ ). This range in  $EF$ s can be attributed to difference in fuel type and burning conditions  
393 (smoldering vs. flaming). We also observed, however, a difference in the  $EF_{\text{group}}$  of dicarboxylic  
394 acid between two tropical peats from the same geographical area (Malaysian peat:  $EF = 35.33 \pm 2.9$   
395  $\text{mg kg}^{-1}$  and Malaysian agricultural peat:  $12.29 \pm 1.02 \text{ mg kg}^{-1}$ ). The highest  $EF$  for individual  
396 dicarboxylic acids was observed for azelaic acid. For example, for Pskov peat the  $EF$  was  $32.1 \pm 4.1$   
397  $\text{mg kg}^{-1}$ ; and for Moscow peat it, was  $17.6 \pm 2.6 \text{ mg kg}^{-1}$ . Azelaic acids were mostly found in the  
398 particulate phase (Fig. 2b, Table S1) and their relative abundances in the gas phase varied between  
399 0.77% (for Pskov peat) and 2.85% (for Eucalyptus). Maleic acid was mostly found in the gas phase  
400 (73%–83%), since it is a lower MW compound ( $MW = 116.0 \text{ g mol}^{-1}$ ) compared to azelaic  
401 ( $MW = 188.22 \text{ g mol}^{-1}$ ) and adipic ( $146.14 \text{ g mol}^{-1}$ ) acids. Succinic and methyl-succinic acids are  
402 found in both gas and particulate phases (Table S1), and their abundance in the particulate phase  
403 was 19–59% and 53–100%, respectively. For Malaysian peat BB emissions, succinic acid was  
404 present only in the particulate phase. A distinguishable increase in dicarboxylic acid mass  
405 concentrations was observed for ambient aerosols followed by a biomass burning event (Cao et  
406 al., 2017) compared to normal ambient concentrations. The formation of saturated dicarboxylic  
407 acids (e.g., succinic acid) and unsaturated dicarboxylic acids (e.g., maleic acid) also was reported  
408 for ambient aerosols collected near a biomass-burning event (Graham et al., 2002; Kundu et al.,  
409 2010; Zhu et al., 2018) and in ice core records historically affected by biomass burning (Müller-  
410 Tautges et al., 2016).





412

413 Figure 2c. EFs for monocarboxylic acids in both particulate phase (solid bars, filter samples) and

414 gas-phase (open bars, XAD samples) from fresh biomass burning emissions for six different fuel

415 types. We did not burn fuels in replicates and standard deviations (SD) were calculated based on

416 replicate analysis of emissions from similar fuels (with identical experimental conditions) during

417 our previous combustion campaigns (Yatavelli et al., 2017) where SD ranged between 9.4 and

418 12% for monocarboxylic acids.

419

420 Monocarboxylic acids can constitute up to 30–40% of total identified organic aerosol mass from  
421 BB emissions (Oros et al., 2006). In our study, we characterized the range from C<sub>6</sub>–C<sub>24</sub>, where  
422 some unsaturated monocarboxylic acids (e.g., oleic acid) also are included. For Alaskan and  
423 Malaysian peat fresh emissions (Fig. 2c), the highest EF (gas + particle) among all analysed  
424 monocarboxylic acids was for hexadecanoic acid (C<sub>16</sub>) with EFs of 55.7±6.6 mg kg<sup>-1</sup> and 51.8±6.2  
425 mg kg<sup>-1</sup>, respectively. The dominance of hexadecanoic acid among other monocarboxylic acids in  
426 combustion emissions also was observed in ambient measurements during biomass-burning events  
427 in southeast Asia (Fang et al., 1999). For Moscow (41.5±6.5 mg kg<sup>-1</sup>) and Pskov (49.8±7.8 mg kg<sup>-1</sup>)  
428 peats, tetradecanoic acid (C<sub>14</sub>) had the highest EFs in fresh samples (Fig. 2c). For Eucalyptus  
429 and Malaysian agricultural peat fresh samples, the largest contributor to monocarboxylic acids was  
430 tetracosanoic acid (C<sub>24</sub>) (Fig. 2c) with EFs of 3.81±0.5 mg kg<sup>-1</sup> and 42.0±5.1 mg kg<sup>-1</sup>, respectively.  
431 As we expected, low MW monocarboxylic acids like hexanoic acid (MW=116 g mol<sup>-1</sup>) was mostly  
432 present in the gas phase, and the gas phase mass fraction varied between 72% (for Malaysian  
433 agricultural peat) and 98% (for Moscow peat). Similar trends were observed for other **low MW**  
434 monocarboxylic acids. For example, the relative abundance of octanoic acid (C<sub>8</sub>) in the gas phase  
435 was 93.4% for Alaskan peat. **High MW** monocarboxylic acids (C<sub>16</sub>>) abundance in the gas phase  
436 was < 2% for all analyzed fuels.

437  
438 Carbon preference index (CPI) is generally used for source apportionment of organic aerosols  
439 (Fang et al., 1999). We also computed the carbon preference index for monocarboxylic acids from  
440 all analyzed fuel combustion emissions by taking even carbon number over odd carbon number  
441 ratio on EFs of monocarboxylic acids ranges from C<sub>6</sub> to C<sub>24</sub> (Fig. S1). For fresh emission samples,  
442 the CPI values ranged from 1.28 (for Moscow peat) to 4.53 (for Eucalyptus). The CPI values are  
443 higher for fresh emissions of tropical peats (for example, 2.78 for Malaysian peat) than for  
444 emissions from peats from high latitudes (for example, 1.74 for Pskov peat). An average CPI index  
445 of 3.7 for monocarboxylic acids was reported for combustion emissions from sedimentary bogs  
446 (Freimuth et al., 2019), in the range of our reported values (CPI: 1.3–4.5).

447  
448 The sums of the EFs for 25 monocarboxylic acids are shown in Fig. 3c. The EF<sub>group</sub> was in the  
449 range of 5 and 515 mg kg<sup>-1</sup> for all fuels. This range is comparable to the EF<sub>group</sub> reported previously

450 for grass (tundra, cotton, Pampas and ryegrass) combustion (32–250 mg kg<sup>-1</sup>) (Oros et al., 2006).  
451 Overall, the trend of low EFs associated with flaming combustion of Eucalyptus (16±0.7 mg kg<sup>-1</sup>)  
452 compared to smoldering peat combustion is also evident for this compound category. Combustion  
453 of peat fuels from tropical regions (e.g., Malaysian agricultural peat) resulted in monocarboxylic  
454 acids EFs of 212±9.6 mg kg<sup>-1</sup> compared to higher EFs from Alaskan peat combustion (505±23 mg  
455 kg<sup>-1</sup>). The origin of monocarboxylic acid is mostly plant wax and oils (Simoneit, 2002). The  
456 relative proportion of plant wax and oils can vary widely among vegetation taxa and also their  
457 concentrations in peat depend on biogeochemical processes involved in peat formation. The  
458 differences in relative abundance of waxes and plant oils in living vegetation and the differences  
459 between biogeochemical processes involved in peat formation for arctic and tropical regions may  
460 be responsible for diverse EFs for monocarboxylic acids.

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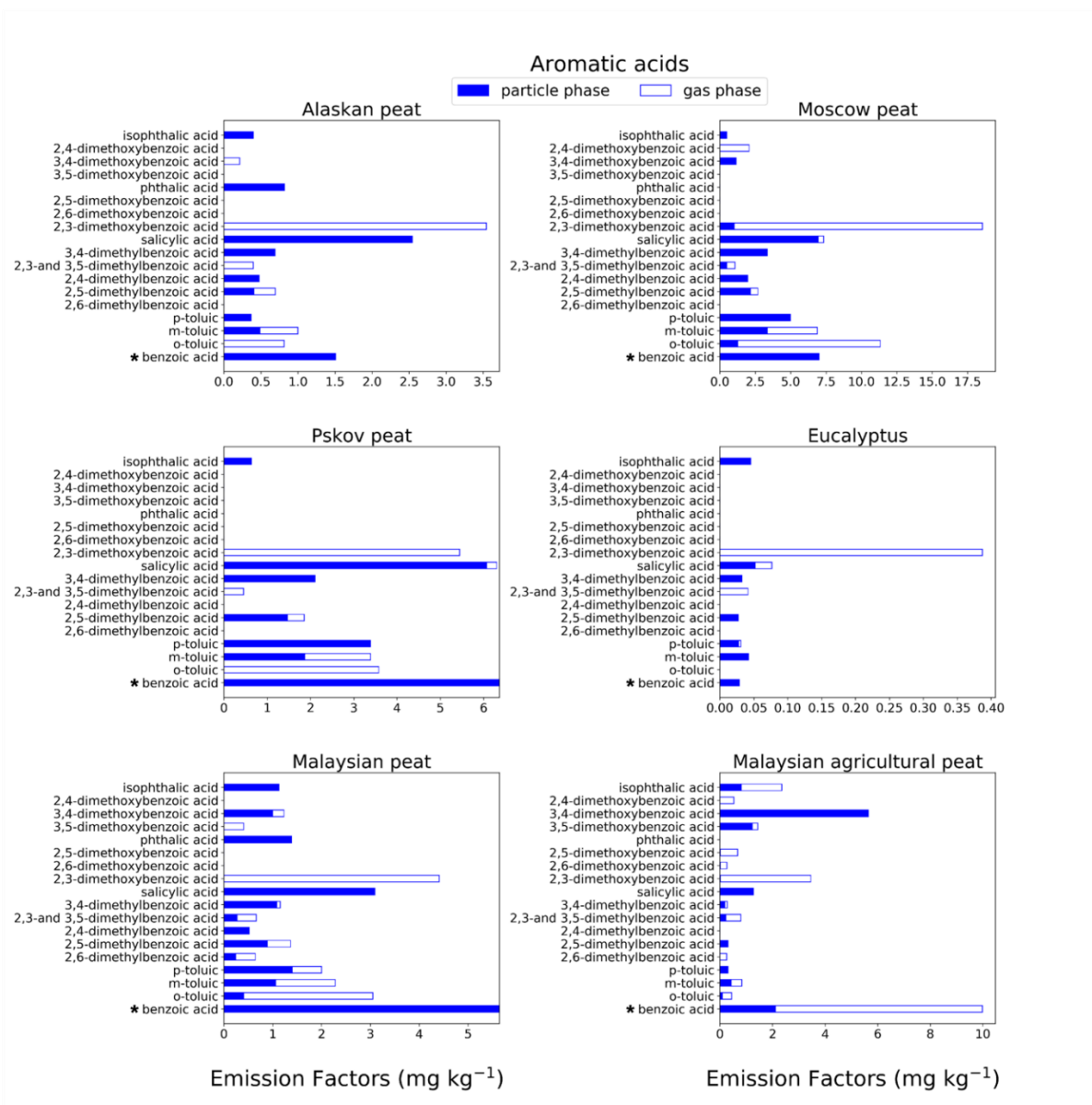
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476  
 477 Figure 2d. EFs for aromatic acids in both particulate phase (solid bars, filter samples) and gas-  
 478 phase (open bars, XAD samples) from fresh biomass burning emissions for six different fuel types.  
 479 We did not burn fuels in replicates, and SD were calculated based on replicate analysis of  
 480 emissions from similar fuels (with identical experimental conditions) during our previous  
 481 combustion campaigns (Yatavelli et al., 2017) where SD ranged between 9.5 and 15% for aromatic  
 482 acids; \*Benzoic acid was found in high concentrations in the XAD blanks that introduced a  
 483 substantial uncertainty to quantification of this compound.

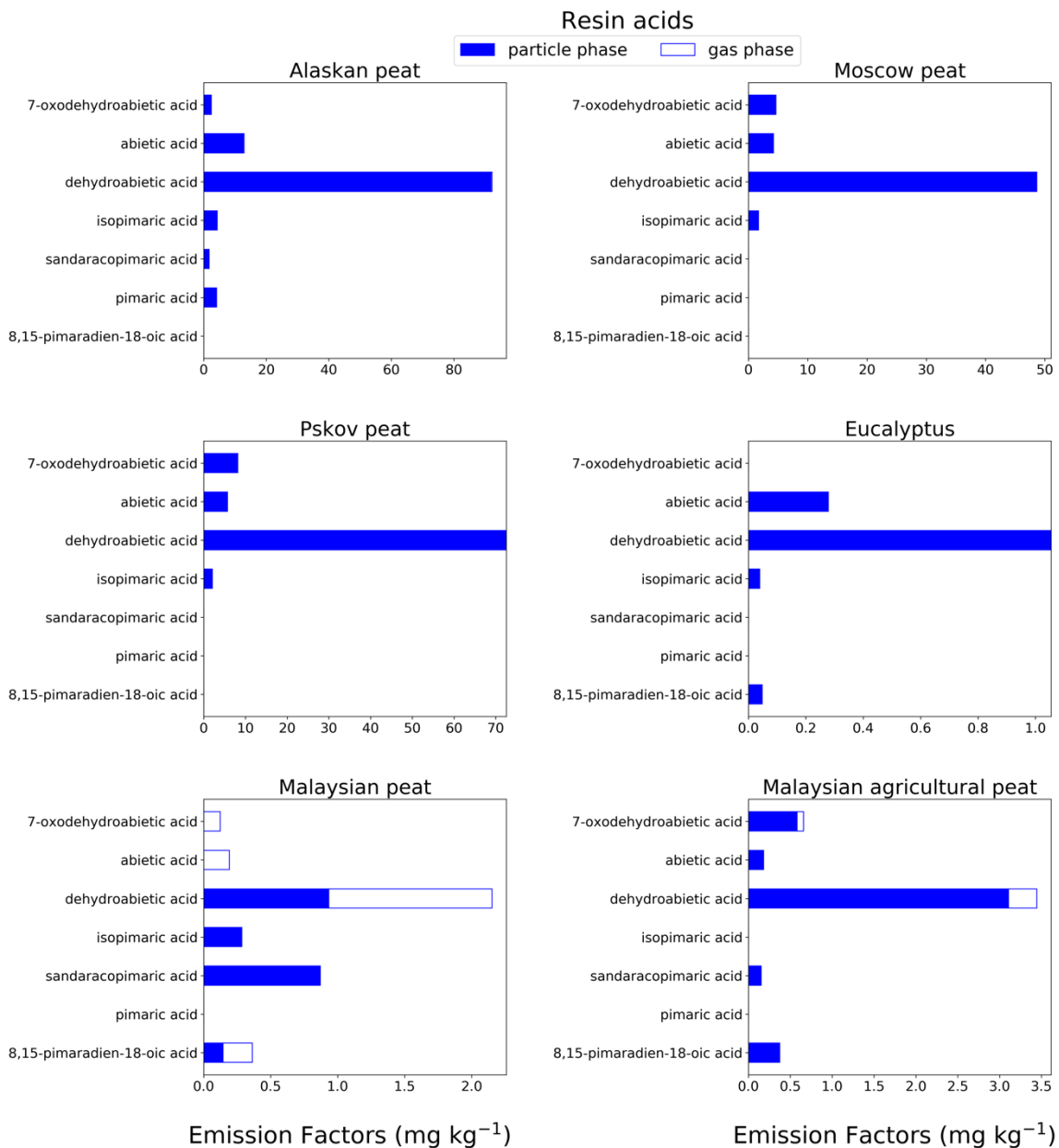
484 Aromatic acids from BB emissions can contribute up to 20–35 % of total identified organic mass  
485 (Wan et al., 2019). In our study, the aromatic acids (e.g., p-hydroxy benzoic acid), excluded  
486 methoxyphenol derivatives and resin acids. For most of the fuels, LMW low MW aromatic acids  
487 (MW <150 g mol<sup>-1</sup>) (e.g., benzoic acid, o-/m-/p-toluic acids) contributed more (almost 40% of  
488 total aromatic acid emissions) toward total fresh emissions, compared to high MW aromatic acids  
489 (MW >150 g mol<sup>-1</sup>). For example, the benzoic acid EF for Malaysian agricultural peat fresh  
490 emission (Fig. 2d) is 9.98±1 mg kg<sup>-1</sup>, and the EF for the same benzoic acid is 6.36±0.6 mg kg<sup>-1</sup> for  
491 Pskov peat (Fig. 2d). Although, o-toluic and p-toluic acids are found in gas phase with 50%–100%  
492 abundance in Alaskan and Moscow peat, benzoic acid is found only in particulate phase. Benzoic  
493 acid was found in high concentrations in the XAD blanks that introduced a substantial uncertainty  
494 to quantification of this compound. One of the most abundant aromatic acids in fresh peat  
495 emissions was 2,3-dimethoxy benzoic acid. For example, for Moscow peat the EF was 18.6±4.7  
496 mg kg<sup>-1</sup>. The acid was mostly found in gas phase (91%–100%) for all fuels (Fig. 2d, Table S1).  
497 2,3-dimethoxy benzoic acid is potentially derived from combustion of lignin moieties of biomass,  
498 and the emission of this compound is more than an order of magnitude lower in emissions from  
499 flaming combustion samples (EF=0.38 ±0.09 mg kg<sup>-1</sup>) than in emissions from smoldering  
500 combustion emissions (EF=5.44±1.36 mg kg<sup>-1</sup>). Emissions of 3,4 dimethoxy benzoic acid were  
501 only observed for peats from tropical regions (EF=5.64±0.8 mg kg<sup>-1</sup>) with 80–100% abundance in  
502 particulate phase. This compound can be used for source apportionment of aerosols emitted from  
503 burning of tropical peat and also can potentially help to distinguish between emissions from  
504 tropical and high latitude peatland fires.

505

506 The EF<sub>group</sub> for aromatic acids in fresh combustion emissions from eucalyptus fuel is extremely  
507 low (0.71±0.05 mg kg<sup>-1</sup>) compared to that for peat fuels (13–69 mg kg<sup>-1</sup>). Among all peat samples,  
508 the Alaskan peat fresh EF was the lowest EF (13.5±0.9 mg kg<sup>-1</sup>), whereas Moscow peat fresh  
509 emissions yielded the highest EF (69±4.4 mg kg<sup>-1</sup>). The difference in total aromatic acid emissions  
510 can be attributed to the variation in the lignin content of the fuels and burning conditions (Simoneit,  
511 2002).

512

513



515

516 Figure 2e. EFs of resin acids in both particulate phase (solid bars, filter samples) and gas-phase  
 517 (open bars, XAD samples) from fresh biomass-burning emissions for six different biomass types.  
 518 As in prior cases, we did not burn fuels in replicates, and SD were calculated based on replicate  
 519 analysis of similar fuels (with same experimental conditions) from our previous campaigns  
 520 (Yatavelli et al., 2017) and SD varies from 9.7–15% for resin acids

521 We quantitatively analyzed combustion emissions for isomers in six resin acids (Table S1). The  
522 most abundant resin acid (78% of total resin acid emission) is dehydroabietic acid (C<sub>20</sub>) that does  
523 not have isomers. The preponderance of this acid over other resin acids in emissions from oak and  
524 pine biomass burning was reported by Simoneit et al. (1993). We found that dehydroabietic acid  
525 (C<sub>20</sub>) content in fresh emissions is 15–30 times higher in fuels from high-latitude peatlands than in  
526 those of tropical origin. For example, the EF for dehydroabietic acid in fresh Alaskan peat  
527 emissions is 92.2±14 mg kg<sup>-1</sup> (Fig. 2e), whereas the same in fresh Malaysian peat emissions is  
528 3.44±0.5 mg kg<sup>-1</sup> (Fig. 2e). Resin acids are supposed to be found mostly in particulate phase based  
529 on their MW and functional groups (Asher et al., 2002; Karlberg et al., 1988; Pankow and Asher,  
530 2008), confirmed by our results (80–100% in particulate phase) with the exception of Malaysian  
531 peat emissions where 56.6% abundance of dehydroabietic acid was found in gas phase. Although  
532 a distinct peak of dehydroabietic acid was observed at the desired retention time for this sample  
533 during GC-MS analysis, we believe this result can be attributed to some unknown interreference  
534 from our analysis procedure.

535  
536 The EF<sub>group</sub> for seven resin acids are presented in Fig. 3e. High EF<sub>group</sub> was observed for Alaskan  
537 (117±15 mg kg<sup>-1</sup>), Pskov (89±12 mg kg<sup>-1</sup>), and Moscow (59±7.7 mg kg<sup>-1</sup>) peats representing mid  
538 latitude and arctic peats. Resin acids (e.g., pimaric acid) are biosynthesized mainly by conifers  
539 (gymnosperms) in temperate regions. In previous work, Iinuma et al. (2007) gave a range of resin  
540 acids EFs from 0 to 110 mg kg<sup>-1</sup>, in agreement with our results. Very low resin acids EFs were  
541 found for peat from tropical regions (e.g., 4 mg kg<sup>-1</sup> for Malaysian peat fresh samples). As  
542 deciduous trees in tropical zones are not prolific resin and mucilage (gum) producers,  
543 compositional data on smoke from such sources should not be expected to show moderate  
544 concentrations of resin acids. This is supported by earlier work by Iinuma et al. (2007), where resin  
545 acids were not even detectable for emissions from Indonesian peat combustion.

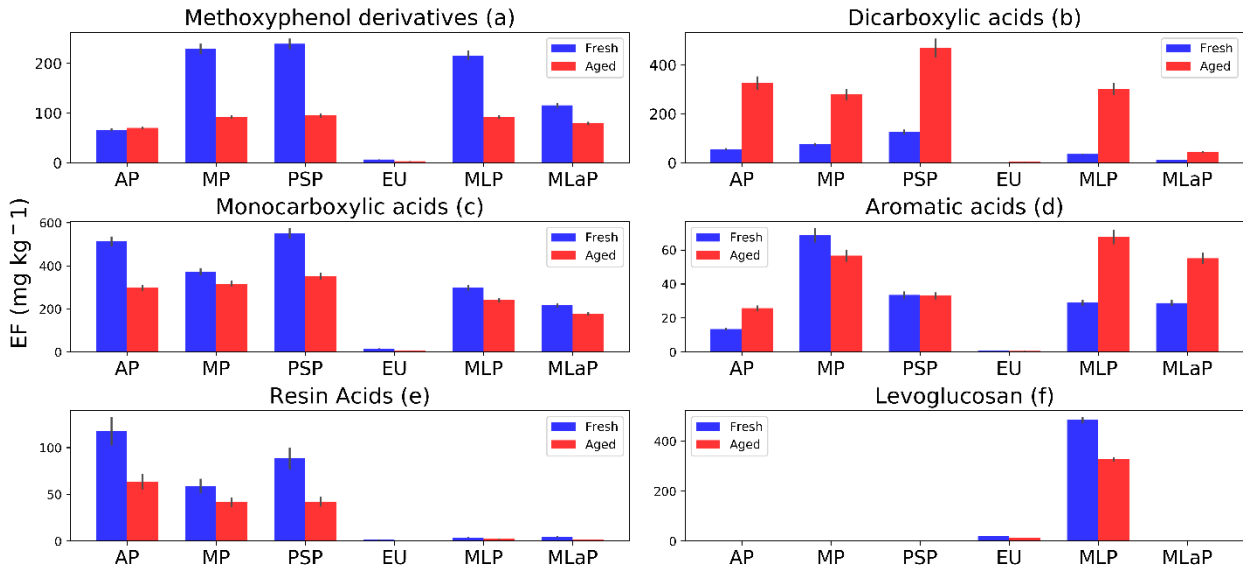
### 546 547 **3.1.5 Levoglucosan**

548 Levoglucosan can be found mostly in particulate phase (Simoneit, 2002). We report levoglucosan  
549 EFs from our analysis of the quartz filter using the IC-PAD technique (no gas phase EFs reported).

550 The EFs of levoglucosan (Fig. 3f) were found to be  $20.9 \pm 0.68 \text{ mg kg}^{-1}$  and  $485 \pm 11.8 \text{ mg kg}^{-1}$  for  
 551 Eucalyptus and Malaysian peat, respectively, and their carbon content is approximately 1.8% and  
 552 2.5 % of the total organic carbon mass characterized by the thermo-optical technique. Fine et al.  
 553 (2002) reported 9% to 16% contribution of levoglucosan to total OC from residential wood  
 554 combustion, a relatively higher percentage than values obtained in our study. Anhydrosugars (e.g.,  
 555 levoglucosan and its isomers) are found in great abundance and have been widely used as a BB  
 556 tracer because of their atmospheric stability, as summarized by Bhattarai et al. (2019). We found  
 557 that levoglucosan constituted 36% and 51% of GC-MS characterized polar (listed in our method)  
 558 organic aerosol mass for eucalyptus and Malaysian peat, respectively, which also is consistent with  
 559 the previous BB literature assembled in the recent review article by Bhattarai et al. (2019).

560  
 561 **3.2. Emission factors of total (gas + particle) organic compounds of six chemical groups and**  
 562 **their changes upon OFR oxidations**

563  
 564 Here we describe changes in the  $EF_{\text{group}}$  followed by OFR oxidation for all six chemical groups.  
 565 Levoglucosan and the most abundant resin acid, dehydroabietic acid also are reported in this  
 566 section.



567  
 568 **Figure 3.** Fuel-based emission factors (EFs) of organic compounds assigned to six chemical  
 569 groups for six different fuels: Alaskan peat (AP), Moscow peat (MP), Pskov peat (PSP),



570 Eucalyptus (EU), Malaysian peat (MLP), Malaysian agricultural peat (MLaP). EFs are presented  
571 as a sum of gas- and particle-phase species mass measured in fresh and OFR-aged BB emissions  
572 in units of  $\text{mg kg}^{-1}$  (mass of emissions per fuel mass combusted). We did not burn fuels in  
573 replicates, and standard deviations (SD) of all chemical groups were calculated based on replicate  
574 analysis of similar fuels (with same experimental conditions) from our previous campaign based  
575 on the data reported by Yatavelli et al. (2017).

576

### 577 *3.2.1. Methoxyphenol derivatives after OFR oxidation (Fig. 3a)*

578 Methoxyphenols can undergo gas-phase oxidation reactions via either aromatic ring  
579 fragmentation/opening to form short-chain ketones, acids, esters, and double bonds in conjugations  
580 with all functional groups or further hydroxylation of aromatic rings to form multiple substituted  
581 aromatic compounds (Yee et al., 2013). In either case, a decrease in methoxyphenols after  
582 oxidation was expected. In our study, a decrease in methoxyphenol's  $\text{EF}_{\text{group}}$  with OFR oxidations  
583 were observed for all fuels (e.g., for Pskov peat from  $239 \pm 11 \text{ mg kg}^{-1}$  to  $95 \pm 4 \text{ mg kg}^{-1}$ ) except for  
584 Alaskan peat, where an insignificant increase from  $66 \pm 3 \text{ mg kg}^{-1}$  to  $70 \pm 3 \text{ mg kg}^{-1}$  after the OFR  
585 oxidation was observed.

586

### 587 *3.2.2. Dicarboxylic acid group after OFR oxidation (Fig. 3b)*

588 A significant increase (2.5–8.5 times) in the  $\text{EF}_{\text{group}}$  of dicarboxylic acids was observed for OFR-  
589 aged samples. For example, the  $\text{EF}_{\text{group}}$  of dicarboxylic acids increased from  $35 \pm 3 \text{ mg kg}^{-1}$  to  
590  $301 \pm 25 \text{ mg kg}^{-1}$  for Malaysian peat and from  $56 \pm 5 \text{ mg kg}^{-1}$  to  $326 \pm 27 \text{ mg kg}^{-1}$  for Alaskan peat.  
591 Oxidation of aerosols potentially produces more oxygenated functional groups (Jimenez et al.,  
592 2009), demonstrated by an increase in O:C ratios (from 0.45 to 0.65) in recent laboratory oxidation  
593 of BB emissions by Bertrand et al. (2018), where TAG-AMS was used to identify the fate of  
594 organic compounds. In this work, however, the number of identifiable compounds with highly  
595 functional groups is constrained by the elution technique used in the TAG method. Our results on  
596 the fate of BB organic aerosols with 18 dicarboxylic acids can provide better mechanistic  
597 understanding about the processes inside OFR.

598

599 3.2.3. Monocarboxylic acid group after OFR oxidation (Fig. 3c)

600 We observed a decrease in monocarboxylic acids  $EF_{\text{group}}$  from OFR aging for all fuels. For  
601 example, the  $EF_{\text{group}}$  for monocarboxylic acids from Alaskan peat combustion decreased from  
602  $514 \pm 23 \text{ mg kg}^{-1}$  (fresh) to  $298 \pm 14 \text{ mg kg}^{-1}$  (aged). A relatively small decrease compared to Alaskan  
603 peat was observed for Malaysian agricultural peat (from  $216 \pm 10 \text{ mg kg}^{-1}$  [fresh] to  $179 \pm 8 \text{ mg kg}^{-1}$   
604 [aged]) and Malaysian peat (from  $298 \pm 14 \text{ mg kg}^{-1}$  [fresh] to  $240 \pm 11 \text{ mg kg}^{-1}$  [aged]) too. This is  
605 probably because of the formation of low MW monocarboxylic acids (e.g., hexanoic acids;  
606  $MW=116 \text{ g mol}^{-1}$ ) after OFR oxidation demonstrated in Fig. 4c and Table S2c and will be  
607 discussed further in section 3.3. Monocarboxylic acids can be oxidized in the atmosphere  
608 (Charbouillot et al., 2012), leading to the formation of dicarboxylic acids from  $C_2$  to  $C_6$  (Ervens et  
609 al., 2004). This is consistent with our results (Figs. 3b, 3c, Table S2b) Moreover, monocarboxylic  
610 acids, during their atmospheric transformations, can produce a potential precursor for formation  
611 of high MW compounds, such as Humic Like Substance (HULIS) (Carlton et al., 2007; Tan et al.,  
612 2012).

613

614 3.2.4. Aromatic acid group after OFR oxidation (Fig. 3d)

615 Levels of aromatic acids from peat burning increased for Alaskan peat, Malaysian peat, and  
616 Malaysian agricultural peat (e.g., from  $29 \pm 2 \text{ mg kg}^{-1}$  to  $68 \pm 4 \text{ mg kg}^{-1}$  for Malaysian peat) by OFR  
617 aging (Fig. 3e, Table S2e). This increase could be from oxidation of phenols and methoxyphenols  
618 in the OFR chamber (Akagi et al., 2011; Legrand et al., 2016). For eucalyptus, Moscow and Pskov  
619 peats it was an insignificantly small decrease (e.g., from  $69 \pm 4 \text{ mg kg}^{-1}$  to  $57 \pm 3 \text{ mg kg}^{-1}$  for  
620 Moscow peat and from  $34 \pm 2 \text{ mg kg}^{-1}$  to  $33 \pm 2 \text{ mg kg}^{-1}$  for Pskov peat). This small decrease in the  
621  $EF_{\text{group}}$  of monocarboxylic acids is statistically insignificant. The oxidation processes occurring in  
622 the OFR are complex, especially in the case of multi-component BB emissions. The decrease  
623 observed for aromatic acids after the OFR may be attributed, however, to multiple generations of  
624 oxidation leading to the breaking of aromatic rings and formation of low MW organic compounds  
625 via fragmentation (Jimenez et al., 2009).

626

627

628 3.2.5. *Resin acids after OFR oxidation (Figs. 3d and 4e)*

629 Resin acids can be oxidized to corresponding oxo-acids (e.g., 7-oxodehydroabietic acid (Karlberg  
630 et al., 1988)), and they are considered to be stronger contact allergens than the resin acids  
631 themselves (Sadhra et al., 1998). Our data showed a small decrease in 7-oxodehydroabietic acid  
632 levels after the OFR (e.g., from  $8.2 \pm 0.8$  to  $4.8 \pm 0.5$  mg kg<sup>-1</sup> for Pskov peat). We noted a significant  
633 decrease in the EF<sub>group</sub> of resin acids from  $117 \pm 15$  mg kg<sup>-1</sup> (fresh) and  $63 \pm 8$  mg kg<sup>-1</sup> (aged) for  
634 Alaskan peat after OFR oxidation, mostly because of some individual compounds like  
635 dehydroabietic acid. Although resin acids are considered to be stable atmospheric tracers for  
636 biomass burning (Simoneit et al., 1993a), we observed a decrease in dehydroabietic acid (most  
637 abundant) EF after the OFR-oxidation of emissions from all fuels (Fig. S2b). For example, for  
638 Alaskan peat (Fig. S2b), the decrease was from  $92 \pm 14$  mg kg<sup>-1</sup> to  $57 \pm 8$  mg kg<sup>-1</sup> after OFR-  
639 oxidation. The fate of resin acids during OFR aging, however, was beyond the scope of this work  
640 and may be the subject of future investigations.

641

642 3.2.6. *Levoglucosan after OFR oxidation (Fig. 3f)*

643 Levoglucosan is one of the most popular tracers of BB emissions, since it has been considered a  
644 stable compound in the atmosphere (Oros et al., 2006; Simoneit, 2002; Simoneit et al., 1999).  
645 Several laboratory studies, however, have demonstrated degradation of levoglucosan in the  
646 presence of OH radicals (Hennigan et al., 2010). Here we observed a decrease of 30% in  
647 levoglucosan levels following OFR oxidation. For example, Malaysian peat decreased from  
648  $485 \pm 12$  mg kg<sup>-1</sup> to  $327 \pm 8$  mg kg<sup>-1</sup>. For eucalyptus, the decrease was from  $20 \pm 0.7$  mg kg<sup>-1</sup> to  $14 \pm 0.6$   
649 mg kg<sup>-1</sup>. This decrease also can be attributed to the degradation process during OH oxidation  
650 (Hoffmann et al., 2010). Levoglucosan oxidation should be studied more, so it can be adequately  
651 used as a tracer of BB emissions.

652

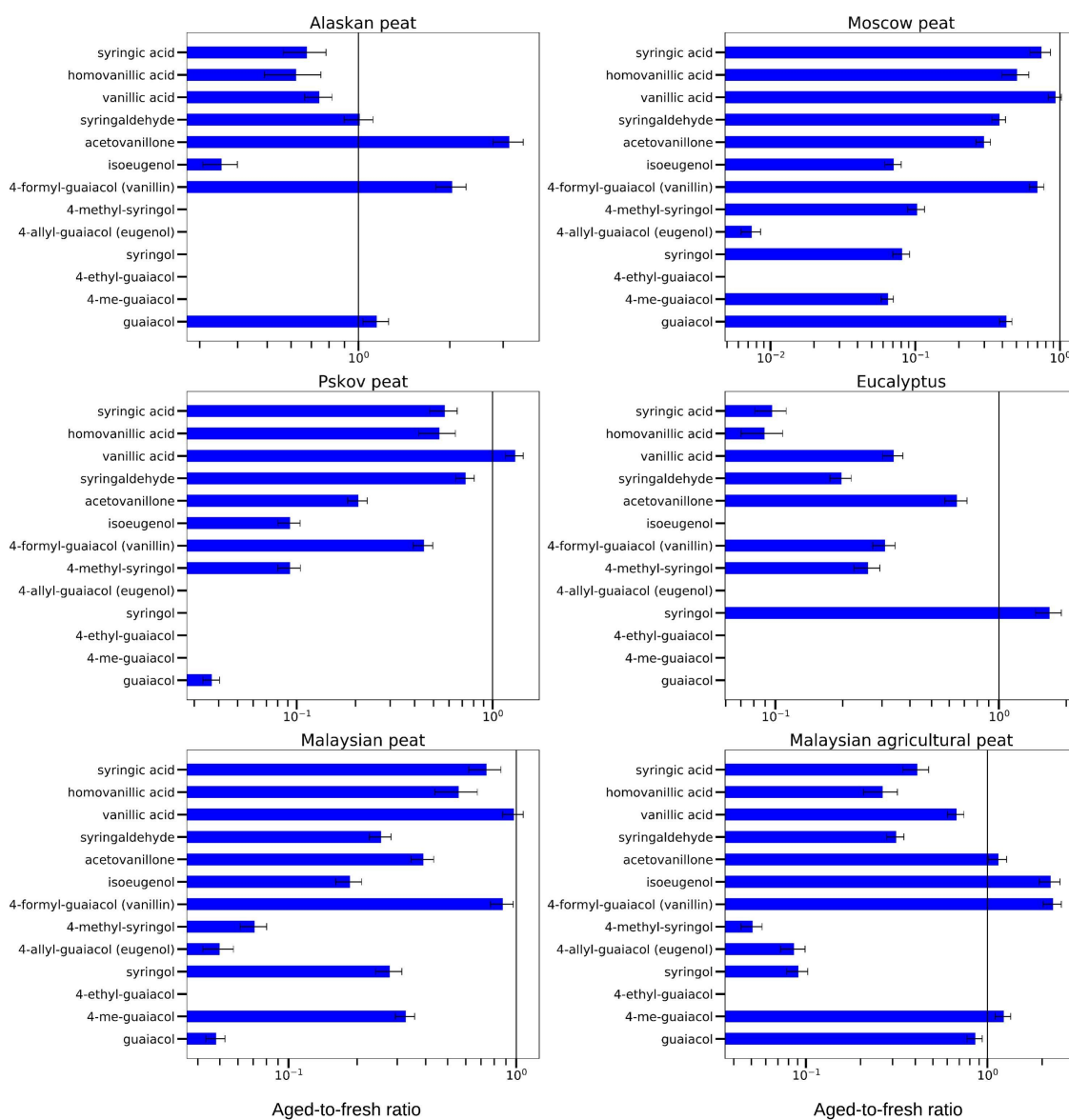
653 **3.3. Aged-to-fresh ratios of total (gas + particle) emission factors of individual organic**  
 654 **compounds assigned to six chemical groups and their changes upon OFR oxidation**

655 We computed aged-to-fresh ratios of individual compounds for all fuels. If the aged-to-fresh ratio  
 656 of one compound is greater than one, this implies that the compound is formed during OFR  
 657 oxidations; if the ratio is less than one, then the compound must have decomposed inside the OFR.

658

659 **3.3.1 Methoxyphenol Derivatives**

Methoxyphenol derivatives

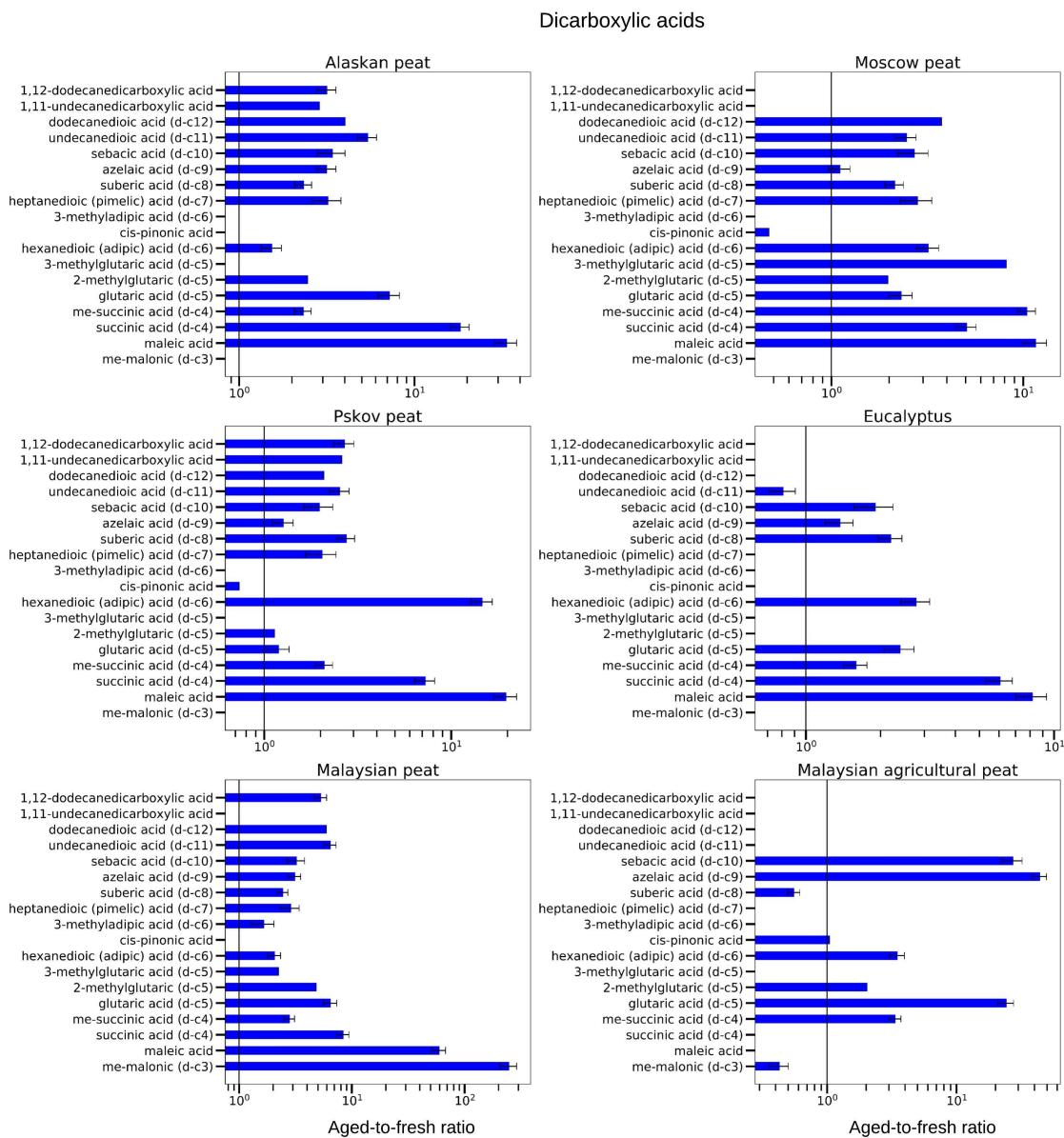


660

661 **Figure 4a.** Aged-to-fresh ratios of total (gas + particle) EFs for methoxyphenols from biomass  
662 burning emissions for six different biomass types presented in log scale. We did not burn fuels in  
663 replicates, and standard deviations (SD) were calculated based on replicate analysis of similar fuels  
664 (with same experimental conditions) from our previous campaigns. SD values derived from EFs  
665 were scaled to ratio.

666  
667 Overall, we found that abundances for methoxyphenol derivatives rapidly decreased upon OFR-  
668 oxidation (Fig. 4a, Table S2a). Some compounds—vanillic acid, acetovanillone, and syringic  
669 acids— demonstrated both increasing and decreasing trends. For example, for Pskov peat, the  
670 aged-to-fresh ratio of guaiacol was  $0.04 \pm 0.01$  reflecting a significant decrease during OFR  
671 oxidation. For Pskov peat, we also observed a ratio less than one for vanillin ( $0.44 \pm 0.05$ ),  
672 indicating vanillin also decreased during OFR oxidation for the same fuel but not to the extent of  
673 guaiacol. At the same time and for the same fuel, a slight increase (aged-to-fresh ratio  $>1$ ) in  
674 vanillic acid was observed ( $1.30 \pm 0.13$ ) in the OFR-oxidized sample. This increase in vanillic acid  
675 concentration can be attributed to the oxidation of vanillin, one of the abundant methoxyphenol in  
676 the fresh emissions from Pskov peat (Fig. 4a, Table S2a). For combustion of other peats, vanillic  
677 acid concentrations also decreased (e.g., aged-to-fresh ratios were  $0.74 \pm 0.08$  and  $0.67 \pm 0.07$  for  
678 Alaskan peat and Malaysian agricultural peat, respectively). Acetovanillone increased by a factor  
679 of three during OFR oxidation for Alaskan peat and around 15 % for Malaysian agricultural peat  
680 (aged-to-fresh ratio  $1.15 \pm 0.13$ ), but the increase for Malaysian agricultural peat was not  
681 **statistically** significant. For other fuels, acetovanillone decreased during the OFR oxidation. For  
682 example, for Moscow Peat, the aged-to-fresh ratio for acetovanillone was  $0.30 \pm 0.03$ . We still need  
683 to investigate the reason why both acetovanillone and vanillic acid increased for some fuels and  
684 decreased for others. The reduction of acetovanillone and vanillic acid was because of a photo-  
685 chemical decomposition process in the OFR with formation of lower **MW** products, such as  
686 succinic acid and maleic acid (Schnitzler and Abbatt, 2018).

687



689

690 **Figure 4b.** Aged-to-fresh ratios of total (gas + particle) EFs for dicarboxylic acids from biomass  
 691 burning emissions for six different biomass types presented in log scale. We did not burn fuels in  
 692 replicates, and SD were calculated based on replicate analysis of similar fuels (with same  
 693 experimental conditions) from our previous campaigns. SD values derived from EFs were scaled  
 694 to ratio.

695

696 In the case of dicarboxylic acids, we observed 2–20 times increase in EFs, but the degree of  
697 enhancement of **low MW** dicarboxylic acids EFs was higher than for **high MW** dicarboxylic acids  
698 EFs. For example, a 20-fold increase in maleic acid, a **low MW** dicarboxylic acid (MW= 116.07  
699 g mol<sup>-1</sup>), was observed during OFR oxidation of Pskov peat emissions (aged-to-fresh ratio =  
700 19.6±2.8), whereas 1,11-undecanedicarboxylic acid, a **high MW** dicarboxylic acid (MW = 244.33  
701 g mol<sup>-1</sup>), EF increased 2.6 times (aged-to-fresh ratio = **2.60±0.01**) for the same fuel. Similarly, the  
702 concentration of succinic acid, a **low MW** dicarboxylic acid (MW= 118.09 g mol<sup>-1</sup>), increased  
703 almost by five times after OFR oxidation (aged-to-fresh ratio = 5.07±0.62), whereas that of  
704 undecanedioic acid, a **high MW** dicarboxylic acid (MW = 230.30 g mol<sup>-1</sup>), increased 2.5 times  
705 (aged-to-fresh ratio = 2.46±0.30) for Moscow peat. This trend was in accord with results from  
706 ambient observations after BB events (Cao et al., 2017; Kawamura and Bikkina, 2016).

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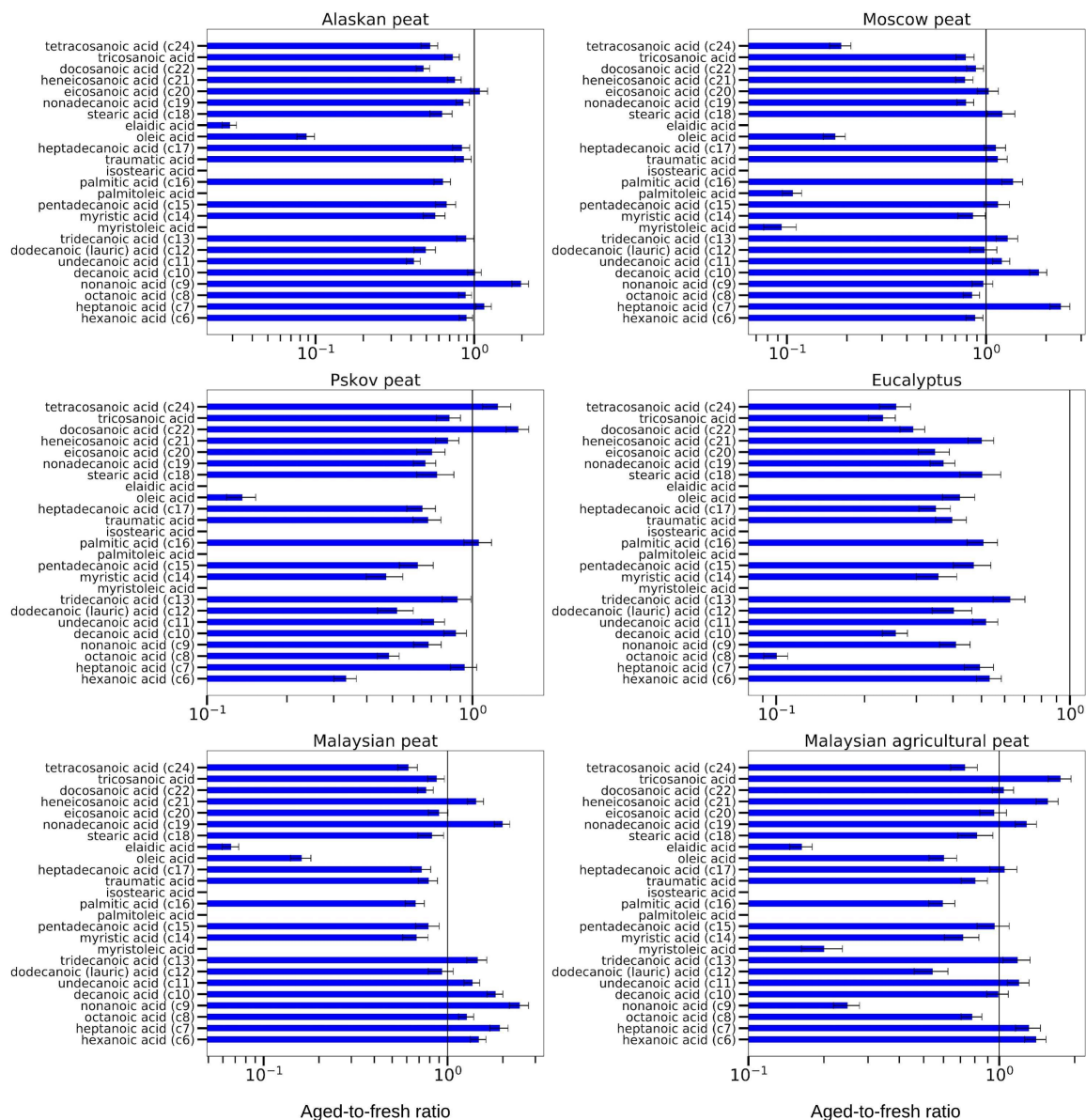
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Monocarboxylic acids



721  
 722 **Figure 4c.** Aged-to-fresh ratios of total (gas + particle) EFs for monocarboxylic acids from  
 723 biomass burning emissions for six different biomass types presented in log scale. We did not burn  
 724 fuels in replicates, and SD were calculated based on replicate analysis of similar fuels (with same  
 725 experimental conditions) from our previous campaigns. SD values derived from EFs were scaled  
 726 to ratio.

727



728 Our analysis of OFR-aged samples showed that concentrations of monocarboxylic acids with  
729 different **MWs** changed during OFR oxidation, but the changes varied from one fuel to another.  
730 For example, the EF of hexanoic acid ( $C_6$ ) was reduced for Eucalyptus (aged-to-fresh ratio =  
731  $0.5\pm 0.05$ ) and fuels from other high-latitude peatlands like Alaskan (aged-to-fresh ratio =  
732  $0.89\pm 0.09$ ), Moscow (aged-to-fresh ratio =  $0.88\pm 0.09$ ) and Pskov peat (aged-to-fresh ratio =  
733  $0.33\pm 0.03$ ). The reduction of hexadecenoic acid was **statistically** insignificant for Alaskan and  
734 Moscow peat. The peats from tropical regions showed exactly the opposite change. Hexanoic acid  
735 increased for both Malaysian (aged-to-fresh ratio =  $1.47\pm 0.14$ ) and Malaysian agricultural peat  
736 (aged- to-fresh ratio =  $1.40\pm 0.14$ ). We observed a similar trend for heptanoic acid ( $C_7$ ) during OFR  
737 oxidation. The tropical peats clearly demonstrated increases (for example, aged-to-fresh ratio =  
738  $1.92\pm 0.22$  for Malaysian peat) in heptanoic acid concentration. For Moscow peat, even though  
739 hexanoic acid concentrations were insignificantly decreased, heptanoic acid concentrations  
740 increased significantly (aged-to-fresh ratio =  $2.37\pm 0.27$ ). This contrast between changes in  
741 hexanoic and heptanoic acid can be explained by a decrease in CPI indices during OFR oxidations  
742 (for example, from 2.78 to 1.7 for Malaysian peat). The reduction of CPI indices indicated that  
743 during oxidation more monocarboxylic acids with odd carbon numbers were formed than  
744 monocarboxylic acids with even carbon numbers. The abundance of hexadecenoic acid ( $C_{16}$ ) was  
745 reduced during OFR oxidation for all fuels (for example, aged-to-fresh ratio =  $0.63\pm 0.08$  for  
746 Alaskan peat) except for Moscow and Pskov peat aged-to-fresh ratio =  $1.05\pm 0.13$  for peat), and  
747 we believe this small increase is **statistically** insignificant. Similarly, tetracosanoic acid ( $C_{20}$ ) was  
748 reduced for all fuels (for example, aged-to-fresh ratio =  $0.61\pm 0.07$  for Malaysian peat) except for  
749 Pskov peat (aged-to-fresh ratio =  $1.24\pm 0.15$ ). The small increase in tetracosanoic acid ( $C_{20}$ )  
750 concentration was again **statistically** insignificant. Even though our results indicated the possibility  
751 of fragmentation of **high MW** monocarboxylic acids and formation of **low MW** monocarboxylic  
752 acids, this was because of the complexity of the OFR oxidation environment. We are not able to  
753 hypothesize what is the main reactive mechanism.

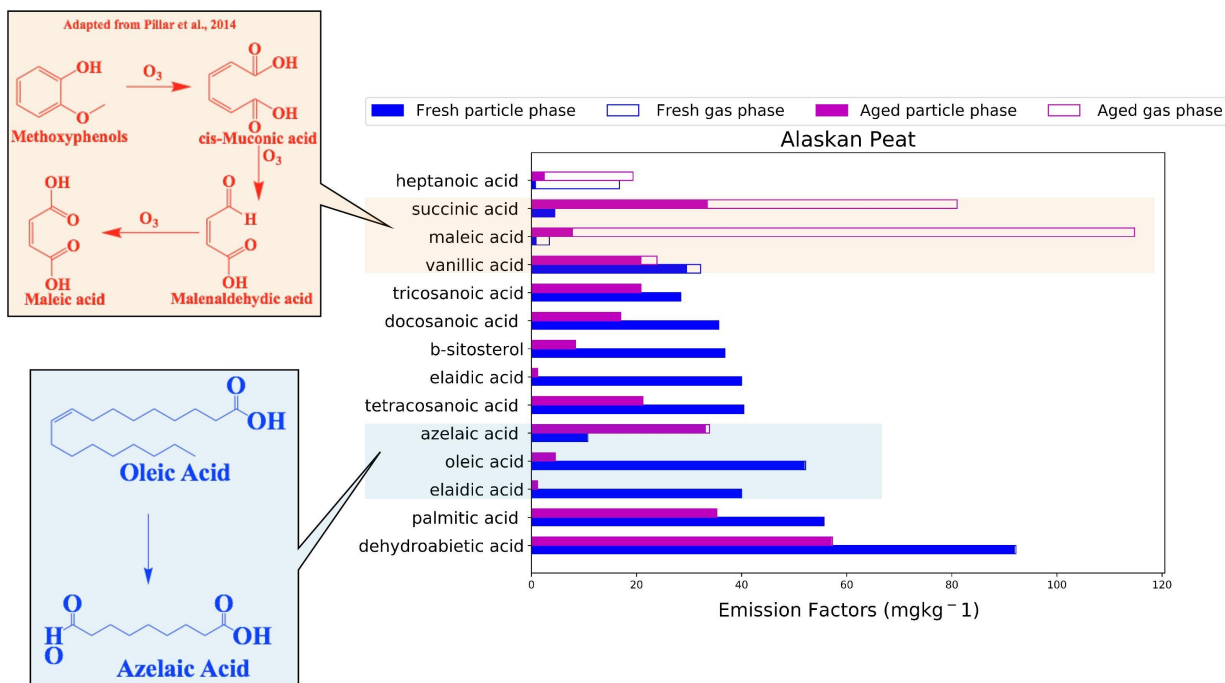
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758 **3.4. Most Contributing Compounds**



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761 **Figure 5.** Emission factors (EF) of top contributing organic compounds assigned to Alaskan peat.

762 Top 10 contributing EFs compounds were selected from both fresh and aged emissions based on

763 EF values. Since fresh and aged samples do not have the same set of compounds after the selection,

764 we included the top 10 compounds for both fresh and aged emissions. Hence, the number of top

765 contributing compounds varies from one fuel to another. Solid bars of each type represent a

766 chemical group from particulate emission of BB fuels and open bars of the same color represent

767 gas phase BB emissions.

768

769 We have identified top contributing compounds for both fresh and aged BB emissions of each fuel

770 to understand how emissions vary from one fuel to another. The top 10 compounds for fresh and

771 aged emission were different, and we merged the top 10 compounds from fresh and aged emissions

772 resulting in different numbers of total top compounds for different fuels. Here we discuss Alaskan

773 peat emissions in their particulate phase (with solid bars) and gas phase (with open bars) as an

774 example, while the remaining results are given in the SM (Fig. S3 and Fig. S4). It is clear that the

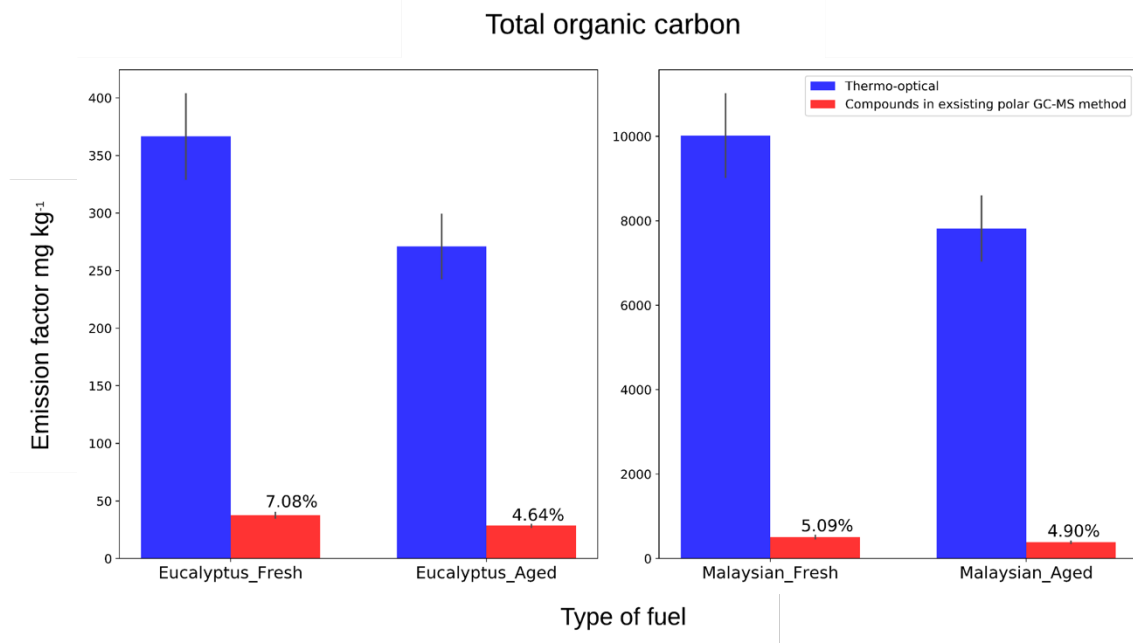
775 top compounds vary between fuels, likely because of both different chemical nature of these fuels  
776 and nature of combustion.

777  
778 Dehydroabietic acid, a resin acid, is the compound with the highest EF (92.2 mg kg<sup>-1</sup>) for fresh  
779 combustion emissions from Alaskan peat. Moncarboxylic acids including palmitic acid (EF = 55.7  
780 mg kg<sup>-1</sup>), tetracosanoic acid (EF = 40.35 mg kg<sup>-1</sup>), and docosanoic acid (EF = 35.38 mg kg<sup>-1</sup>) were  
781 also found in high abundance in fresh emissions from the combustion of this fuel. The high  
782 contributions of  $\beta$ -sitosterol (EF = 36.84 mg kg<sup>-1</sup>) and alkenoic acids (e.g., oleic acid EF = 52.1  
783 mg kg<sup>-1</sup>) to emissions are unique to Alaskan peat. All the compounds described above are found  
784 in particulate phase. After the OFR-oxidation, both dehydroabietic acid and  $\beta$ -sitosterol, considered  
785 to be potential markers for biomass burning emissions (Simoneit et al., 1993b), decreased from  
786 91.9 mg kg<sup>-1</sup> to 57.2 mg kg<sup>-1</sup> and 36.8 mg kg<sup>-1</sup> to 8.38 mg kg<sup>-1</sup> in particulate phase, respectively.  
787 This reduction in EF because of OFR oxidation for both dehydroabietic acid and  $\beta$ -sitosterol must  
788 be considered when using these compounds as biomass-burning markers. We observed the  
789 formation of low MW organic compounds, particularly in gas phase, from OFR oxidation. For  
790 example, the EF of heptanoic acid increased from 2.42 mg kg<sup>-1</sup> to 16.9 mg kg<sup>-1</sup> and that of maleic  
791 acid increased from 7.8 mg kg<sup>-1</sup> to 107 mg kg<sup>-1</sup> in the gas phase because of OFR oxidation. Such  
792 a significant increase in the EF of maleic acid can be explained by the aqueous phase oxidation of  
793 methoxyphenols (El Zein et al., 2015) in the presence of ozone. We found that the oxidation inside  
794 the chamber was happening under dry conditions and understand that the reactions of organic  
795 compounds with OH radicals inside the OFR chamber will prevail over reactions with ozone. As  
796 we had our ozone scrubbers placed after sampling media (Fig. 1) to prevent the pumps and online  
797 instruments from ozone-induced damage, we suspect that the maleic acid was not formed inside  
798 the OFR chamber but rather by potential oxidation of organic compounds on filters with relatively  
799 longer exposure of ozone (40–60 min for smoldering combustion). Succinic acid EFs increased in  
800 both the gas phase (from 0.0 mg kg<sup>-1</sup> to 47.5 mg kg<sup>-1</sup>) and the particulate phase (from 4.43 mg kg<sup>-1</sup>  
801 to 33.5 mg kg<sup>-1</sup>). Azelaic acid EFs showed mainly an increase in the particulate phase (from 10.5  
802 mg kg<sup>-1</sup> to 33.1 mg kg<sup>-1</sup>), and we think that this was because of the oxidation of oleic and eladic  
803 acid during OFR oxidation.

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806 **3.5 Contribution of polar fraction to total organic carbon**  
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808  
809 **Figure 6.** Contribution of GC-MS characterized polar compound carbon mass to total thermo-  
810 optical organic carbon mass. The y-axis shows the total carbon mass with dimensions of mass per  
811 mass. The error bars represent analytical uncertainties of the methods. For the thermo-optical  
812 method, uncertainties are the standard deviation of results from multiple punches on the same filter  
813 and for the GC-MS method, uncertainties were computed by taking the square root of sum of the  
814 squares of individual analytical uncertainties of all compounds included.

815  
816 For Figure 6, we calculated the carbon content of total GC-MS characterized mass of identified  
817 polar organic compounds and compared results with the total OC mass characterized by the  
818 thermo-optical technique to estimate the contribution of polar compounds. The OC emissions were  
819 higher for smoldering combustion ( $10,209 \pm 5 \text{ mg kg}^{-1}$  for Malaysian peat fresh emissions) than  
820 for flaming combustion ( $366.5 \pm 7 \text{ mg kg}^{-1}$  for eucalyptus fresh emissions) samples, similar to the  
821 observation of flaming and smoldering combustion by Akagi et al. (2011). Total OC emissions  
822 are highly dependent on the type of fuel. For example, the fuel-based OC emission factor for rice  
823 crop residue burning is  $1960 \text{ mg kg}^{-1}$  (Cao et al., 2008), whereas burning of corn and conifer forest

824 yields emission factors of  $\sim 1457 \text{ mg kg}^{-1}$  (Andreae and Rosenfeld, 2008) and  $\sim 7800 \text{ mg kg}^{-1}$   
825 (Akagi et al., 2011), respectively. Figure 6 shows that the 84 identified polar compounds in our  
826 study constituted 4.5% to 7% of total OC mass for both fresh and aged emissions. From Indonesian  
827 peat combustion emissions, Jayarathne et al. (2018) were able to identify polar compounds that  
828 constituted 5.446% of total organic carbon mass. In recent work, based on both a field campaign  
829 with prescribed burning and laboratory investigations, Jen et al. (2019) quantified a fraction (10–  
830 65%) of only identified compounds (not a fraction of total mass) by the use of the 2D-GC-MS  
831 technique. In our work, we identified only up to 7% (Fig. 6) of the total particle phase OC, and  
832 further analysis of unidentified compounds is needed to improve understanding of atmospheric  
833 chemistry of BB emissions.

834

835

#### 836 **4. Summary and conclusions**

837 In this study, we chemically characterized the polar fraction of biomass-burning aerosols from  
838 laboratory combustion of six different globally and regionally important fuels—five of them  
839 representing smoldering and one of them representing flaming combustion. Our objective was to  
840 understand how emissions of the polar compounds (e.g., methoxyphenols) varied from one fuel to  
841 another during these combustion experiments and what are the relative distribution of these polar  
842 compounds in gas and particulate phase. We also identified the fates of these polar compounds  
843 following laboratory oxidation/aging (OFR aging). Resin acids were found mostly in emissions  
844 from combustion of peats from high latitude regions but not in emissions of tropical peatlands  
845 (e.g.,  $EF_{\text{group}} = 117 \pm 15 \text{ mg kg}^{-1}$  for Alaskan peat and  $EF_{\text{group}} = 4.0 \pm 0.5 \text{ mg kg}^{-1}$  for Malaysian peat).  
846 Similarly, monocarboxylic acids were found in higher abundance in emissions from high latitude  
847 peatlands compared to tropical peatland emissions (e.g.,  $EF_{\text{group}} = 505 \pm 36 \text{ mg kg}^{-1}$  for Alaskan peat  
848 and  $EF_{\text{group}} = 212 \pm 15 \text{ mg kg}^{-1}$  for Malaysian agricultural peat). The presence of both guaiacol and  
849 syringol moieties in all fuels indicated a part of the biomass, considered as representative of a  
850 particular geographical region, is deciduous for all fuels. **Low MW** compounds are mostly found  
851 in gas phase (e.g., guaiacol found in gas phase 82–100%), whereas high **MW** (e.g., **high molecular**  
852 **weight**—monocarboxylic acids [ $>C_{16}$ ] more than 98% for all fuels) and highly oxygenated  
853 compounds (e.g., syringic acid and acetovanillone 65–100% in particulate phase) are found in

854 particulate phase with high abundance. Monocarboxylic acids (1.2–3 times) and methoxyphenols  
855 (1.5–2.5 times) decreased after OFR oxidation, whereas dicarboxylic acids increased by 3–9 times  
856 followed by OFR oxidation. Relatively **low MW** hexanoic acid increased for both Malaysian  
857 (aged-to-fresh ratio =  $1.47 \pm 0.14$ ) and Malaysian agricultural peat (aged- to-fresh ratio =  
858  $1.40 \pm 0.14$ ), whereas **high MW** tetracosanoic acid ( $C_{20}$ ) was reduced for all fuel increases (e.g.,  
859 aged-to-fresh ratio =  $0.61 \pm 0.07$  for Malaysian peat). This indicated fragmentation occurring inside  
860 the OFR chamber. With relative distribution of the top 10–15 compounds from Alaskan peat, we  
861 were able to identify transformation of unsaturated fatty acids (e.g., oleic acid) to dicarboxylic  
862 acids (e.g., azelaic acid). We identified only up to 7% of the total particle phase OC and further  
863 analysis of unidentified compounds with GC-MS full scan is needed for better understanding of  
864 atmospheric chemistry of BB emissions.

865

866

867 **Data availability.** Data can be provided upon request: <andrey.khlystov@dri.edu>.

868

869 **Author contributions.** DS, VS, and AK designed experiments. **DS and CB performed sample,**  
870 **data collection, and extraction.** DS performed **extractions,** derivatizations, GC-MS analysis,  
871 summarized data, and wrote the paper. AW provided biomass fuels. VS, AK, and HM provided  
872 input on interpretation of results. VS, HM, and AK revised the manuscript.

873

874 **Competing interests.** The authors declare that they have no conflict of interest.

875

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882 associated with high maleic acid formation. The authors also thank Rodger Kreidberg for revising  
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884

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