

Response to interactive comment on “Significant production of ClNO₂ and possible source of Cl₂ from N₂O₅ uptake at a suburban site in eastern China” by Men Xia et al. from anonymous Referee #1

The manuscript "Significant production of ClNO₂ and possible source of Cl₂ from N₂O₅ uptake at a suburban site in eastern China" by Xia et al. presents a set of measurements of nitryl chloride (ClNO₂) and molecular chlorine (Cl₂) taken near the city of Nanjing, in Eastern China, in April 2018. The authors use this dataset, and related observations, to analyze the formation of ClNO₂ and Cl₂ and to draw conclusions about the underlying multiphase chemical mechanism. The paper is well written and the data are presented in a clear and concise way. The analysis and the results are sound and the authors propose some novel ideas that will certainly be of great interest to the community. I only have a few, fairly minor, comments, but overall I think this paper is suitable for publication in Atmospheric Chemistry and Physics.

Response: we appreciate the reviewer for the positive comments and helpful suggestions. Below is the response to each comment. The reviewers' comments are italicized followed by our responses and changes shown in blue and red, respectively. And the corrections are also marked as red color in the revised manuscript. Please note that the line numbers mentioned below refer to the original submission (line numbers in the revised version have changed).

General Comments —————

In Section 3.3, the authors discuss the calculation of the yield of ClNO₂, comparing the "BT" parametrization by Bertram and Thornton (2009) with a new parametrization. Looking at figure 4, I am not sure I completely agree with the author's interpretation. The new parametrization proposed in this paper does indeed agree better with the observations for yields between 0.4 and 0.6; however I would argue that the agreement is worse than the BT parametrization at higher yields (around 0.8) and only slightly better at lower yields (below 0.4). Clearly, the relationship between the various parameters is more complicated than either parametrization assume, and perhaps this suggests that there are other parameters that are not currently taken into account which play a role. In any case, I suggest that the authors revise their statements in this section (and the related parts of the conclusions and the abstract) to be more accurate.

Response: we appreciate and agree your comment on our interpretation of the performance of the new parameterization at higher yields (0.75~1). We also agree that other unconstrained factors may influence $\phi(\text{ClNO}_2)$. We have revised the relevant texts as below.

Revision in the main text:

Line 329-330 (section 3.3): The parameterized $\phi(\text{ClNO}_2)_{\text{BT+Org}}$ better matches the observed $\phi(\text{ClNO}_2)$ at low to median yields (0~0.75) and the R² and slope values in the linear regression are closer to 1 (Fig. 4b). However, the parameterized $\phi(\text{ClNO}_2)_{\text{BT+Org}}$ is smaller than the observed $\phi(\text{ClNO}_2)$ at high yields (0.75~0.9), which may be attributable to other unconstrained factors in the parameterization, e.g., mixing state and phase state issues.

In Section 3.4, the authors propose a mechanism for the production of Cl₂ during the night.

The key point of the argument is that, for the observations to be consistent with each other, $\gamma(\text{ClNO}_2)$ must decrease and there is not really a good explanation for why that would be the case. Although I agree with this logic, there may be other parameters that influence $\gamma(\text{ClNO}_2)$ besides Cl^- , H^+ and D_p . In particular organics, which are mentioned as important for $\gamma(\text{N}_2\text{O}_5)$ in the previous section may inhibit the uptake of ClNO_2 as well. Likewise, RH, other aerosol components, and perhaps even temperature, may have an effect. I appreciate that it is not possible to exhaust all possible parameters but I think the authors should expand their analysis a little bit here, to make a more robust case.

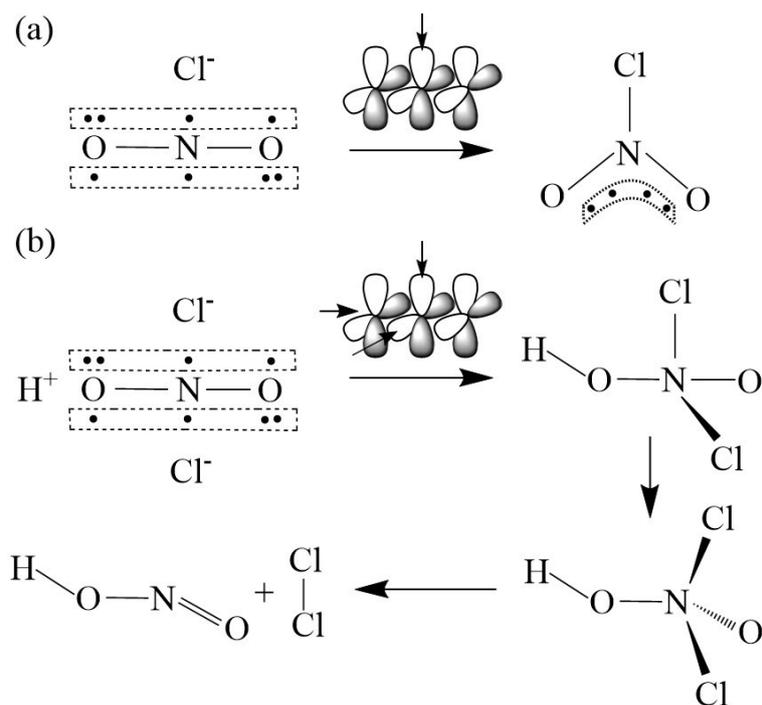
Response: we agree with the referee that other unconstrained factors, in addition to those examined, may influence the $\gamma(\text{ClNO}_2)$. We have now examined the dependence of $\gamma(\text{ClNO}_2)$ on RH, T, and other relevant aerosol components (e.g., NO_3^- , SO_4^{2-} , NH_4^+ , and aerosol organics). Results show no obvious dependence of $\gamma(\text{ClNO}_2)$ on those parameters. We have clarified this point as follows.

Revision in the main text:

Line 398-400 (section 3.4.1): In our study, the D_p was derived from the ratio of wet V_a to S_a by assuming volume-limited uptake (Ammann et al., 2013). We also calculated D_p assuming surface-limited uptake, and no correlation with $\gamma(\text{ClNO}_2)_{\text{obs}}$ was indicated. Moreover, the $\gamma(\text{ClNO}_2)_{\text{obs}}$ showed no obvious relationship with other factors such as T, RH, aerosol liquid water content (ALW), NO_3^- , SO_4^{2-} , NH_4^+ , and aerosol organics (figure not shown).

The authors propose that Cl_2 formation is a co-product of ClNO_2 when N_2O_5 is hydrolyzed on an acidic particle. I would like to see a bit more discussion of this potential mechanism. For ClNO_2 the mechanism is quite straightforward: NO_2^+ reacts with Cl^- to form ClNO_2 . For Cl_2 it does not seem so obvious to me how exactly NO_2^+ and Cl^- interact to form Cl_2 . If the authors have a mechanism in mind please explain or add the relevant reference(s). Otherwise, if this is simply an hypothesis, then please state so clearly.

Response: we agree it would make the contention much more convincing if we can suggest the potential formation mechanism for Cl_2 from reaction of Cl^- and NO_2^+ . Here is our proposed mechanism (see below figure). According to the hybrid orbital theory, the NO_2^+ ion has two non-bonded π molecular orbitals due to participation of the d orbital of the central nitrogen atom (Baird et al., 1981). When Cl^- attacks one of the π molecular orbitals, ClNO_2 is formed. In the same way, Cl^- can attach to the other π molecular orbitals of NO_2^+ and form a short-lived HNO_2Cl_2 intermediate in presence of H^+ . Then, HNO_2Cl_2 decomposes to produce HONO and Cl_2 .



Revision in the main text:

Line 402-404 (section 3.4.1): The mechanism is depicted in Figure 7 and goes as follows. It is known that N_2O_5 hydrolysis on aerosol is responsible for the production of NO_2^+ . According to the hybrid orbital theory, the NO_2^+ ion has two non-bonded π molecular orbitals due to participation of the d orbital of the central nitrogen atom (Baird and Talyer, 1981). ClNO_2 is formed via the nucleophilic addition of Cl^- to one of the π molecular orbitals of NO_2^+ (Figure 7a) (Taylor, 1990; Behnke et al., 1997). In the same way, we propose a side reaction that the second Cl^- can attach to the other π molecular orbital of NO_2^+ and form a short-lived HNO_2Cl_2 intermediate in presence of H^+ . It is proposed that the unstable HNO_2Cl_2 decomposes to produce Cl_2 (and HONO) (Figure 7b). This mechanism can explain concurrent productions of Cl_2 and ClNO_2 from N_2O_5 hydrolysis but needs confirmation by additional laboratory and theoretical studies.

Minor Comments —————

Section 2.1: Are there other relevant parameters (e.g., NO_x) that you can use to compare the two sampling sites?

Response: it is a pity that only simultaneous measurements of O_3 were conducted at both sites.

Section 2.2: Can you please add the detection limits to the text? It would also be useful to see examples of spectra for N_2O_5 , ClNO_2 , Cl_2 and HOCl (these could go in the Supplementary Information).

Response: agreed. We have added the detection limits of N_2O_5 , ClNO_2 , and Cl_2 in the main text, and an example of spectra in the supplementary information. Below is the revision.

Revision in the main text:

Line 153-154 (section 2.2): The detection limits (3σ) of N_2O_5 , ClNO_2 , Cl_2 were 7 pptv, 2 pptv,

and 5 pptv, respectively.

Revision in the SI:

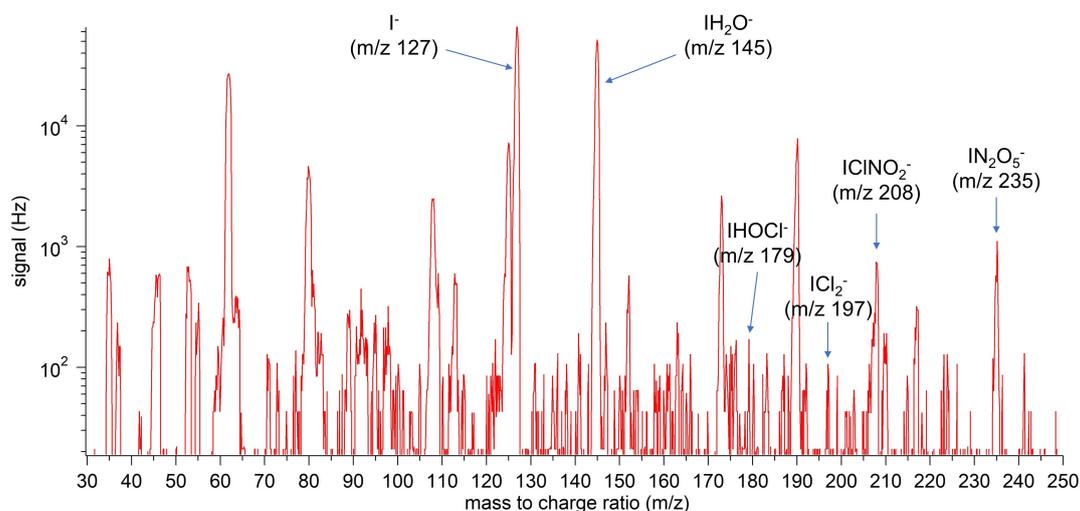


Figure S3. An example of the CIMS spectra taken at 18 April 01:00 LT.

Line 187: what about NO₃ photolysis?

Response: thanks for the reminder of NO₃ photolysis, but in the present study, we focus on the nighttime chemistry of NO₃. So, the $k(NO_3)$ here is the loss rate for nighttime, and photolysis is not included.

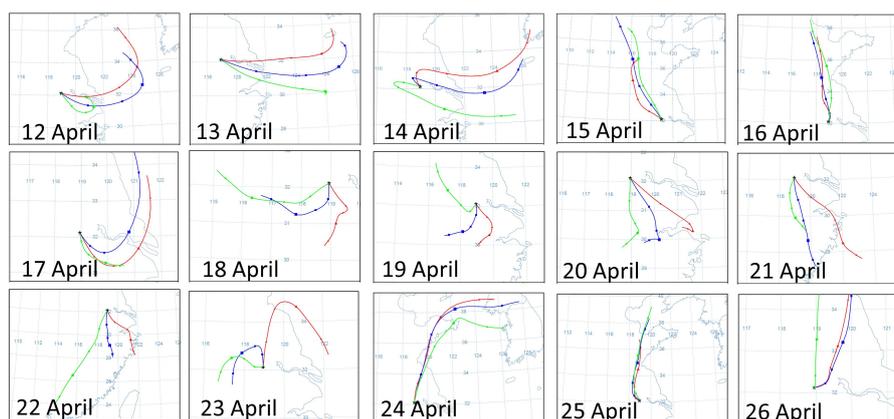
Section 3.2: It seems to me, from figure 3, that the levels of VOC also play a role, not just O₃, RH and Temperature.

Response: the role of VOCs had been included in the calculation the NO₃ reactivity which is dependent on VOC levels. For example, In the plume 3, the NO₃ reactivity due to VOCs decreased compared that in the plume 1. So, a larger proportion of NO₃ was lost via N₂O₅ uptake in the plume 3, which promoted ClNO₂ formation.

Lines 416-423: What about the outflow from Nanjing, which is west of the sampling site? I would think there are industrialized areas also on that part of the country not just between Nanjing and the ocean. Are SO₂ and NO_x very different in the two cases shown in figure 7? Can you please add some detail.

Response: 1. this is a good point. We have examined backward trajectories for the whole observation period but did not identify air masses from urban Nanjing in the west. Please see the figure below. We have added the trajectories figure in the SI.

Revision in the SI:



Red: 200m height. Blue: 500 m height. Green: 1000 m height. Local time: 00:00.

Figure S2. Daily backward trajectories arriving at the sampling sites during the field observation period.

2. The levels of NO_x and SO_2 are slightly higher in marine air compared with continental air in the two cases in figure 7 (see the table below). We have added this point in the main text.

Date	NO_x (ppb)	SO_2 (ppb)	Note
13-Apr	13.1 ± 3.1	3.9 ± 0.1	Marine air passing YRD industry
18-Apr	11.5 ± 0.6	3.3 ± 0.3	Continental air

Revision in the main text:

Line 420-422 (section 3.4.2): The average concentrations of SO_2 (3.9 ± 0.1 ppbv) and NO_x (13.1 ± 3.1 ppbv) in the marine air masses were higher than those (NO_x : 11.5 ± 0.6 ppbv, SO_2 : 3.3 ± 0.3 ppbv) in the inland air masses.

Response to interactive comment on “Significant production of ClNO_2 and possible source of Cl_2 from N_2O_5 uptake at a suburban site in eastern China” by Men Xia et al. from anonymous Referee #2

General Comment

The paper entitled with “Significant production of ClNO_2 and possible source of Cl_2 from N_2O_5 uptake at a suburban site in eastern China” presented comprehensive observations of N_2O_5 , ClNO_2 and Cl_2 as well as other supporting parameters at a regional site in Nanjing. The authors performed a detailed studies on the heterogeneous processes subjected to N_2O_5 uptake and the chlorine productions. Some insights are given on the multiphase chemistry production of Cl_2 . This study further extends the current exploration of the nighttime chemistry in China from North China Plain and Pearl River Delta to Yangtze River Delta which are certainly valuable to be published in ACP. Nevertheless, I think the current analysis needs some further careful check especially for the Section 3.4 as suggested in the follows.

Response: we appreciate the reviewer for the positive comments and helpful suggestions. Below is the response to each comment. The reviewers’ comments are italicized, followed by

our responses and changes shown in blue and red, respectively. And the corrections are also marked as red color in the revised manuscript. Please note that the line numbers mentioned below refer to the original submission (line numbers in the revised version have changed).

Specific Comment

1. Line 149 – 150. More details need to be given for the sentence “the permeation rate of Cl₂ was quantified by chemical titration and ultraviolet spectrophotometry.” How much Cl₂ is generated for calibration and what is the accuracy?

Response: we have added more details of Cl₂ calibration. The permeation rate of Cl₂ generated for calibration is 380 ± 20 ng/min. We have added further details of Cl₂ calibration in the SI. Below is the revision.

Revision in the SI:

Text S1.4. Details of Cl₂ calibration

The Cl₂ standard was generated from a permeation tube heated to 40 °C and flushed by ultrapure nitrogen gas (20 sccm) and then diluted in humidified zero air (6 SLPM). During the field campaign, Cl₂ from the permeation tube was introduced into a KI solution (2% wt) for 1 hour. The permeation rate of Cl₂ (380 ± 20 ng/min) was calculated from the I₃⁻ concentration in the KI solution which was measured by ultraviolet–visible spectrophotometry at 351 nm.

2. Section 3.2. The high ClNO₂ case is of high interest. It would be nice if the authors can try to analyze why the ClNO₂ production become higher for plume 3 than plume 1. The Cl⁻ ion concentrations seem to be quite small and constant for the whole period.

Response: we did analyze the reason for the higher ClNO₂ production in the plume 3. As shown in lines 273-277 and Fig. 3, larger proportion of NO₃ was lost via N₂O₅ uptake in plume 3, which caused elevated ClNO₂ production in plume 3 compared with plume 1. As the Cl⁻ concentration was similar in plumes 1 and 3 (0.17 ± 0.02 and 0.19 ± 0.03 μg/m³, respectively), we did not attribute the higher ClNO₂ in plume 3 to Cl⁻ concentrations.

Low chloride concentrations while high levels of ClNO₂ were also observed in previous studies, where HCl condensation was proposed to replenish particulate chloride to sustain the ClNO₂ production (Osthoff et al., 2008; Thornton et al., 2010). We have clarified this point in the revised manuscript.

Revision in the main text:

Line 277-278 (section 3.2): Compared with the high levels of ClNO₂ (up to 3.5 ppbv) on the night of 17 April, the concentration of Cl⁻ was low and relatively constant (~0.1 ppbv) during that period. The low chloride but high ClNO₂ levels were also observed in previous studies, and HCl partition was proposed to replenish particulate chloride to sustain the ClNO₂ production (Osthoff et al., 2008; Thornton et al., 2010).

3. Line 306-307. “The $\phi(\text{ClNO}_2)$ value ranged from 0.28 to 0.89 (mean, 0.56 ± 0.15), which was among the highest values in the world (McDuffie et al., 2018b).” I suggest to delete “which was among the highest values in the world (McDuffie et al., 2018b).” The $\phi(\text{ClNO}_2)$ is varied within 0-1 depending on the ratio of $[\text{Cl}^-]/[\text{H}_2\text{O}]$, so I do not think the highest is meaningful.

Response: we agree and will delete “which was among the highest values in the world”.

4. Line 338. The equation 11 and corresponding text. I think the estimation and the use of [org] needs more discussion. If the reaction between org and NO₂⁺ is the key to formulate the equation, then the org should be the part of water soluble organics. And I wonder why the reaction with acetate can be similar to the field observations presented herein. What are the major water soluble organics here in this study? And actually you have two adjustable parameters, one is k₅ and the other is the exact [org].

Response: we agree that ideally the [org] here should be water-soluble organics. However, the water-soluble organics are not available in our study, and only total organics were measured on-site. So operationally we assume that organics are all water-soluble, similar to previous studies (McDuffie et al., 2018a; McDuffie et al., 2018b).

The k₅/k₃ value derived here (2.06) happens to be similar to that of acetate. One possibility is that the aerosol organics have a weighted average k₅/k₃ value of 2.06. Due to limited observation species, we don't know the exact composition of water-soluble organics.

k₅ is the only adjustable parameter here. Since we assumed all observed aerosol organics to be water-soluble, the unknown water-soluble proportion of organics is factored in k₅.

Revision in the main text:

Line 331-334 (section 3.3) Here we assumed that the observed aerosol organics were all water-soluble and reactive toward NO₂⁺, as previous studies did (McDuffie et al., 2018a; McDuffie et al., 2018b). The unknown water-soluble proportion of aerosol organics is factored in k₅.

Line 350-352 (section 3.3): A recent laboratory study (Staudt et al., 2019) derived k₅/k₃ = 3.7 for acetate, which happens to be very similar to our results.

5. Line 375-376. The D_p is derived from the ratio of the wet V_a to S_a. As I understood, the the dry D_p is measured directly from SMPS instrument and the wet D_p can be estimated from empirical GF factor or measurements if available. It may be worth to check two kinds of D_p for your calculations, one is for the surface area concentrations when it is surface limited, and the other is for the volume concentrations when it is limited by volume bulk reactions. The calculation of the Gamma_CINO₂ may be influenced by the choice of the different D_p. A slightly different equation is suggested for your test of the gamma_CINO₂.

$$K_{het} = 1 / (Dp/Dg + 4/(gamma_CINO2*c_CINO2)) * 3 * ALW/Dp$$

Dg, gas diffusion constant

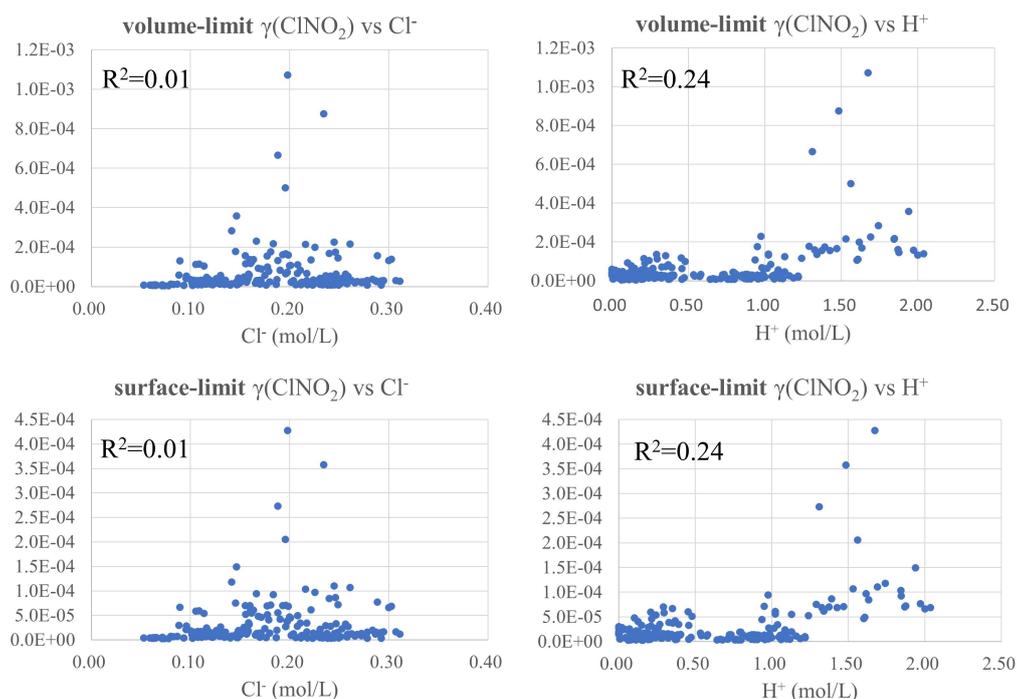
ALW, aerosol liquid water content

Response: we calculated D_p from the V_a/S_a ratio (255.3±32.9 nm in campaign average) based on the volume-limited uptake, we now also calculate D_p according to the average surface area of aerosols, which represents the surface-limited uptake. The surface-limited D_p gives a result of 122.6±26.7 nm in campaign average. The γ(CINO₂) is independent of both the surface-limited D_p and volume-limited D_p.

We appreciate the referee for suggesting a formula to test the γ(CINO₂). After investigating this formula, we suggest that D_p should be replaced with R_p (particle radius) before using this formula. Then, we adopt both volume-limited D_p and surface-limited D_p in this formula and derive γ(CINO₂), respectively. Results are shown in the table below.

Statistics	$\gamma(\text{ClNO}_2)$ in our paper	volume-limited $\gamma(\text{ClNO}_2)$	surface-limited $\gamma(\text{ClNO}_2)$
average value	1.1E-04	6.8E-05	3.1E-05
standard deviation	2.0E-04	1.2E-04	5.1E-05

Then, we investigate the relationship between $\gamma(\text{ClNO}_2)$ (volume-limited and surface-limited, respectively) and Cl^- and H^+ . While no significant correlation is indicated. Results are shown in the figure below.



To sum, we will state in the manuscript that we examine the relationship between $\gamma(\text{ClNO}_2)$ and D_p derived from both volume-limited uptake and surface-limited uptake. Since different methods of calculating $\gamma(\text{ClNO}_2)$ lead to the same conclusion, i.e., no correlation between $\gamma(\text{ClNO}_2)$ and Cl^- and H^+ , we prefer sticking to our original method to derive $\gamma(\text{ClNO}_2)$.

Revision in the main text:

Line 398-400 (section 3.4.1): In our study, the D_p was derived from the ratio of wet V_a to S_a by assuming volume-limited uptake (Ammann et al., 2013). We also calculated D_p assuming surface-limited uptake (diameter of the average surface area), and no correlation with $\gamma(\text{ClNO}_2)_{\text{obs}}$ was indicated.

6. Line 397 – 400. The ALW could be a variable to check for Cl_2 production.

Response: thanks for this suggestion. we have checked ALW in Cl_2 production by investigating the dependence of $\gamma(\text{ClNO}_2)$ on ALW. However, no obvious correlation is found. We have clarified this point.

Revision in the main text:

Line 398-400 (section 3.4.1): ...Moreover, the $\gamma(\text{ClNO}_2)_{\text{obs}}$ showed no obvious relationship with other factors such as T, RH, **aerosol liquid water content (ALW)**, NO_3^- , SO_4^{2-} , NH_4^+ , and

aerosol organics (figure not shown).

7. Section 3.4.1 and 3.4.2, if the essence of Cl₂ production is from ClONO₂(aq) + H⁺ + Cl⁻ → Cl₂(g) Both the production of ClONO₂ uptake and N₂O₅ uptake which can generate ClONO₂(aq) could be the explanation for the Cl₂ production. The authors may then to quantify the ratio of these two channels from the observations. In addition, the HOCl channel can also be assessed. Response: we think both ClONO₂ uptake and direct N₂O₅ uptake can generate ClONO₂(aq) and then produce Cl₂. However, based on ambient measurements, we cannot separate the contribution of the two pathways, because an assumption to derive γ(ClONO₂) was that Cl₂ was all produced by ClONO₂ uptake (Eq. 13, line 369). We will clarify in the revised draft that N₂O₅ uptake and ClONO₂ uptake are indistinguishable in driving Cl₂ production.

Since HOCl and Cl₂ were poorly correlated, we believe that the HOCl channel has minor contributions to Cl₂ production at our site.

Revision regarding this comment is made together with the comment 8.

8. Section 3.4.2, the analysis of φ(ClONO₂) is only meaningful, if the authors can prove the N₂O₅ uptake is the major (i.e. >90%) production channel of the Cl₂.

Response: we think that the referee meant φ(Cl₂) in the above comment. In this paper, we attempt to demonstrate/argue that the three previously proposed reaction pathways could not explain the observed Cl₂ productions at night at our site, and suggest an additional one. We summarize our views as follows.

1. For ClONO₂ → Cl₂, in section 3.4.1, we showed that the γ(ClONO₂) derived from the assumption that Cl₂ was from ClONO₂ uptake didn't have the expected relationships with Cl⁻, H⁺, and D_p, which are the known factors that influence the ClONO₂ uptake. Thus, we argue that it could not be the main pathway for Cl₂ production at our site. We note that for our proposed new Cl₂ pathway: NO₂⁺ + Cl⁻ + H⁺, we do not rule out the ClONO₂ uptake channel, but assume it can produce Cl₂ via NO₂⁺.

2. For HOCl → Cl₂, we think it is a minor Cl₂ production channel, because Cl₂ was not correlated to HOCl but highly correlated with ClONO₂ during most nights. The same logic was adopted in a previous study (Haskins et al., 2019) to support the view of insignificant role of HOCl.

3. For ClONO₂ → Cl₂, ClONO₂ were not measured in our study. According to a recent field study in north China, nocturnal ClONO₂ levels was low (maximum ~15 pptv). So, we assume that ClONO₂ → Cl₂ was not a significant Cl₂ formation pathway, given the γ(ClONO₂) on the order of 10⁻³ (Burkholder et al., 2015).

Revision for comments 7 and 8 in the main text:

Line 348-352 (section 3.4.1)

Our result suggests that Cl₂ was related to ClONO₂, but the HOCl pathway (R5) and coal burning were of minor importance at our site. ClONO₂ was not measured during our study. Recent field measurements at a rural site in northern China reported low ClONO₂ levels at night (maximum ~ 15 pptv) (Breton et al., 2018). We believe that the ClONO₂ levels at our site was also low, and production pathway (R6) was insignificant given low γ(ClONO₂) (~10⁻³) (Haskins

et al., 2019). At our site, the Cl₂/ClNO₂ ratios varied on different nights, which implies that differences exist in the production efficiencies of Cl₂ relative to those of ClNO₂.

Line 381-383 (section 3.4.1)

For example, the box for 18:00–19:00 contains the $\gamma(\text{ClNO}_2)$ estimated at 18:00–19:00 on 11, 12, and 14 April (Fig. 6b–6d, orange lines). Fig. 6b–6d displays the observed Cl₂ levels (blue lines) and the projected trends of Cl₂ levels from Eq. (12), where the grey lines adopted the highest $\gamma(\text{ClNO}_2)$ value, 6.69×10^{-5} observed in the field study of Haskins et al. (2019). During early evening hours (i.e., 18:00–19:00), the $\gamma(\text{ClNO}_2)$ value derived in our study was one to two orders of magnitude higher than those in that study. This result implies that either ClNO₂ uptake was much faster at our site or other pathways were involved in Cl₂ production. We provide evidence below that the latter is likely the case.

Line 406-411 (section 3.4.1)

We propose a new framework to estimate nighttime Cl₂ production by treating Cl₂, ClNO₂, and most nitrate all ultimately originating from N₂O₅ uptake. We assign a production yield to Cl₂ from the N₂O₅ uptake ($\varphi(\text{Cl}_2)$) analogous to the ClNO₂ yield and calculate this metric using Eq. (14):

$$\varphi(\text{Cl}_2) = \frac{d[\text{Cl}_2]/dt}{k(\text{N}_2\text{O}_5)[\text{N}_2\text{O}_5]} \quad (14)$$

The above formulation does not rule out the production of Cl₂ from the ClNO₂ uptake, because such production, if any, is also a result of N₂O₅ uptake and has thus been incorporated in Eq. (14).

Reference:

- Ammann, M., Cox, R. A., Crowley, J. N., Jenkin, M. E., Mellouki, A., Rossi, M. J., Troe, J., and Wallington, T. J.: Evaluated kinetic and photochemical data for atmospheric chemistry: Volume VI – heterogeneous reactions with liquid substrates, *Atmos. Chem. Phys.*, 13, 8045-8228, 10.5194/acp-13-8045-2013, 2013.
- Baird, N. C., and Taylor, K. F.: The stabilizing effect of d orbitals on the central nitrogen atom in nitrogen-oxygen molecules and ions, *Chemical Physics Letters*, 80, 83-86, [https://doi.org/10.1016/0009-2614\(81\)80062-0](https://doi.org/10.1016/0009-2614(81)80062-0), 1981.
- Behnke, W., George, C., Scheer, V., and Zetzsch, C.: Production and decay of ClNO₂ from the reaction of gaseous N₂O₅ with NaCl solution: Bulk and aerosol experiments, *Journal of Geophysical Research: Atmospheres*, 102, 3795-3804, 1997.
- Bertram, T., and Thornton, J.: Toward a general parameterization of N₂O₅ reactivity on aqueous particles: the competing effects of particle liquid water, nitrate and chloride, *Atmospheric Chemistry and Physics*, 9, 8351-8363, 2009.
- Breton, M. L., Hallquist, Å. M., Pathak, R. K., Simpson, D., Wang, Y., Johansson, J., Zheng, J., Yang, Y., Shang, D., and Wang, H.: Chlorine oxidation of VOCs at a semi-rural site in Beijing: significant chlorine liberation from ClNO₂ and subsequent gas-and particle-phase Cl–VOC production, *Atmospheric Chemistry and Physics*, 18, 13013-13030, 2018.

Brown, S. S., Dubé, W. P., Tham, Y. J., Zha, Q., Xue, L., Poon, S., Wang, Z., Blake, D. R., Tsui, W., and Parrish, D. D.: Nighttime chemistry at a high altitude site above Hong Kong, *Journal of Geophysical Research: Atmospheres*, 121, 2457-2475, 2016.

Burkholder, J. B., Sander, S. P., Abbatt, J. P. D., Barker, J. R., Huie, R. E., Kolb, C. E., et al. (2015). Chemical kinetics and photochemical data for use in atmospheric studies: Evaluation number 18. Pasadena, CA: Jet Propulsion Laboratory, National Aeronautics and Space Administration.

Gaston, C. J., and Thornton, J. A.: Reacto-diffusive length of N_2O_5 in aqueous sulfate-and chloride-containing aerosol particles, *The Journal of Physical Chemistry A*, 120, 1039-1045, 2016.

Haskins, J. D., Lee, B. H., Lopez - Hilifiker, F. D., Peng, Q., Jaeglé, L., Reeves, J. M., Schroder, J. C., Campuzano - Jost, P., Fibiger, D., and McDuffie, E. E.: Observational constraints on the formation of Cl_2 from the reactive uptake of $ClNO_2$ on aerosols in the polluted marine boundary layer, *Journal of Geophysical Research: Atmospheres*, 124, 8851-8869, 2019.

McDuffie, E. E., Fibiger, D. L., Dubé, W. P., Lopez - Hilfiker, F., Lee, B. H., Thornton, J. A., Shah, V., Jaeglé, L., Guo, H., and Weber, R. J.: Heterogeneous N_2O_5 uptake during winter: Aircraft measurements during the 2015 WINTER campaign and critical evaluation of current parameterizations, *Journal of Geophysical Research: Atmospheres*, 123, 4345-4372, 2018a.

McDuffie, E. E., Fibiger, D. L., Dubé, W. P., Lopez Hilfiker, F., Lee, B. H., Jaeglé, L., Guo, H., Weber, R. J., Reeves, J. M., and Weinheimer, A. J.: $ClNO_2$ yields from aircraft measurements during the 2015 WINTER campaign and critical evaluation of the current parameterization, *Journal of Geophysical Research: Atmospheres*, 123, 12,994-913,015, 2018b.

Osthoff, H. D., Roberts, J. M., Ravishankara, A. R., Williams, E. J., Lerner, B. M., Sommariva, R., Bates, T. S., Coffman, D., Quinn, P. K., Dibb, J. E., Stark, H., Burkholder, J. B., Talukdar, R. K., Meagher, J., Fehsenfeld, F. C., and Brown, S. S.: High levels of nitryl chloride in the polluted subtropical marine boundary layer, *Nature Geoscience*, 1, 324-328, 10.1038/ngeo177, 2008.

Roberts, J. M., Osthoff, H. D., Brown, S. S., and Ravishankara, A.: N_2O_5 oxidizes chloride to Cl_2 in acidic atmospheric aerosol, *Science*, 321, 1059-1059, 2008.

Taylor, R., *Electrophilic Aromatic Substitution*, John Wiley, New York, 1990.

Thornton, J. A., Kercher, J. P., Riedel, T. P., Wagner, N. L., Cozic, J., Holloway, J. S., Dube, W. P., Wolfe, G. M., Quinn, P. K., Middlebrook, A. M., Alexander, B., and Brown, S. S.: A large atomic chlorine source inferred from mid-continental reactive nitrogen chemistry, *Nature*, 464, 271-274, 10.1038/nature08905, 2010.

Wang, Z., Wang, W., Tham, Y. J., Li, Q., Wang, H., Wen, L., Wang, X., and Wang, T.: Fast heterogeneous N_2O_5 uptake and $ClNO_2$ production in power plant and industrial plumes observed in the nocturnal residual layer over the North China Plain, *Atmospheric Chemistry and Physics*, 17, 12361-12378, 2017.