Formation of Nighttime Sulfuric Acid from the Ozonolysis of Alkenes in Beijing

Yishuo Guo¹, Chao Yan^{1,2,*}, Chang Li¹, Zemin Feng¹, Ying Zhou¹, Zhuohui Lin¹, Lubna Dada², Dominik Stolzenburg², Rujing Yin³, Jenni Kontkanen², Kaspar R. Daellenbach², Juha Kangasluoma^{1,2}, Lei Yao², Biwu Chu², Yonghong Wang², Runlong Cai², Federico Bianchi², Yongchun Liu¹ and Markku Kulmala^{1,2}

¹ Aerosol and Haze Laboratory, Beijing Advanced Innovation Center for Soft Matter Science and Engineering, Beijing University of Chemical Technology, Beijing, China

² Institute for Atmospheric and Earth System Research / Physics, Faculty of Science, University of Helsinki, Finland

³ State Key Joint Laboratory of Environment Simulation and Pollution Control, State Environmental Protection Key Laboratory of Sources and Control of Air Pollution Complex, School of Environment, Tsinghua University, Beijing, China

*Correspondence to: Chao Yan (chao.yan@helsinki.fi)

15 Abstract. Gaseous sulfuric acid (SA) has received a lot of attention for its crucial role in atmospheric new particle formation (NPF), and for this reason, studies until now have mainly focused on daytime SA when most NPF events occur. While daytime SA production is driven by SO₂ oxidation of OH radical from photochemical origin, the formation of SA during night and its potential influence on particle formation remains poorly understood. Here we present evidence for significant nighttime SA production in urban Beijing during winter, yielding concentrations between 1.0 and 3.0×10^6 cm⁻³. We found a high frequency $(\sim 30\%)$ of nighttime SA events, which are defined by the appearance of a distinct SA peak observed between 20:00 and 04:00 20 local time, and with the maximum concentration exceeding 1.0×10^6 cm⁻³. These events mostly occurred during unpolluted nights with low vapor condensation sink. Furthermore, we found that under very clean conditions (visibility > 16.0 km) with abundant ozone (concentration > 2.0×10^{11} cm⁻³, ~ 7 ppb), the overall sink of SA was strongly correlated with the products of O₃, alkenes and SO₂ concentrations, suggesting that the ozonolysis of alkenes played a major role in nighttime SA formation 25 under such conditions. This is in light with previous studies showing that the ozonolysis of alkenes can form OH radical and stabilized Criegee intermediate (sCI), both of which are able to oxidize SO₂ leading to SA formation. However, we also need to point out that there exist additional sources of SA under more polluted condition, which are not investigated in this study.

Moreover, there was a strong correlation between SA concentration and the number concentration of sub-3 nm particles in

both clean and polluted nights. Different from forest environments, where oxidized biogenic vapors are the main driver of

nighttime clustering, our study demonstrates that the formation of nighttime cluster mode particles in urban environments is

30

5

10

1

mainly driven by nighttime SA production.

Keywords: nighttime SA, urban environment, ozonolysis of alkenes, sub-3 nm particles

35 1. Introduction

40

45

Atmospheric aerosol particles have considerable impact on global climate by directly affecting the radiation balance of the earth and by indirectly acting as cloud condensation nuclei (Stocker et al., 2014). The number concentration of these aerosol particles depends to a large extent on the atmospheric new particle formation (NPF), which includes gas-phase nucleation and subsequent growth of newly formed particles. Studies over the past twenty years have shown that the SA is the major gaseous precursor of NPF in most environments inside the continental boundary layer (Lee et al., 2019). Sulfuric acid driven NPF can proceed as SA-water binary nucleation, SA-water-ammonia ternary nucleation (Kirkby et al., 2011), SA-amine-water nucleation (Almeida et al., 2013;Kuerten et al., 2014), SA-organics nucleation (Riccobono et al., 2014), and SA-organics-amonia nucleation (Lehtipalo et al., 2018) and H₂SO₄-H₂O-NH₃-amine nucleation (Myllys et al., 2019;Yu et al., 2012). Both the nucleation rate (J_{nuc}) and the initial growth rate of newly formed particles tends to have a power-law relationship with the SA concentration: $J_{nue} = k \times SA^{\alpha}$, where the activation nucleation is dominant when $\alpha \approx 1$ (Kulmala et al., 2006), the kinetic nucleation is dominant when $\alpha \approx 2$ (Riipinen et al., 2007;Paasonen et al., 2009;Erupe et al., 2010) and the thermodynamic nucleation becomes more crucial when α is larger than 2.5 (Wang et al., 2011).

Due to the importance of SA for NPF, accurate and reliable measurement of SA is of great importance. Up to now, ambient SA concentrations have been reported for many sites (Weber et al., 1997;Weber et al., 1998;Weber et al., 1999;Paasonen et al., 2010;Jokinen et al., 2018;Fiedler et al., 2005;Eisele et al., 2006;Boy et al., 2008;Iida et al., 2008;Wang et al., 2011;Kuerten et al., 2016;Yao et al., 2018;Mauldin et al., 2001;Erupe et al., 2010;Yu et al., 2014). These studies indicate that the concentration level of SA in the atmosphere is closely related to human activities. In general, daytime SA concentration is around 10⁵ cm⁻³ in pristine Antarctica region (Mauldin et al., 2001), 10⁶ cm⁻³ in remote continental, remote marine and forest regions (which are less affected by human activities) and 10⁷ cm⁻³ in urban and rural agricultural lands (which are influenced dominantly by human activities, (Birmili et al., 2003;Fiedler et al., 2005;McMurry et al., 2005;Iida et al., 2008;Erupe et al., 2010;Paasonen et al., 2010;Wang et al., 2011;Yao et al., 2018)). SA generally shows a distinct diurnal pattern correlating with radiation (Lu et al., 2019) with typical concentrations between 10⁶ to 10⁷ cm⁻³ during daytime and 10⁴ to 10⁶ cm⁻³ during nighttime. The seasonal variation of SA is only reported in very few studies, showing higher concentrations during spring and summer than in winter and autumn (Erupe et al., 2010).

The strong connection between SA and NPF has led previous studies to mostly focus on understanding the SA formation in the daytime. However, recent observation on the formation of sub-3 nm particles have shown that these cluster mode particles also exist with high concentration during the night (Junninen et al., 2008;Lehtipalo et al., 2011;Kulmala et al., 2013;Kecorius et al., 2015;Mazon et al., 2016;Yu et al., 2014) and sometimes even nighttime particle nucleation events can be clearly distinguished. In boreal forest environments, nighttime cluster formation can be attributed to highly oxygenated organic molecules (HOMs) (Kammer et al., 2018;Rose et al., 2018). However, the sources of SA and its role in the particle formation

during the nighttime remain largely unresolved, both of which are the focus of this work. In this study, we show noticeable nighttime sulfuric acid sometimes increase of SA during the nighttime in urban Beijing. We further investigate the main sources of SA and demonstrate its role in the nocturnal formation of sub-3nm clusters.

2. Nighttime Sulfuric Acid Formation

70

75

During the daytime, gaseous SA is primarily a photochemical product generated from the oxidation of SO₂ by OH radical, while at nighttime, SA is highly associated with non-photochemical oxidants, most likely the non-photochemical OH radical and stabilized Criegee intermediate (sCI) (Mauldin et al., 2012;Taipale et al., 2014). And that non-photochemical pathway has been shown important in parameterizing the SA formation the boreal forest environment (Dada et al., 2020). The non-photochemical oxidation pathway mainly includes the following five reactions. First, the production of (stabilized) Criegee Intermediates by the ozonolysis of alkenes:

Alkene +
$$O_3 \xrightarrow{\kappa_1} \phi sCI + (1 - \phi)CI + RCHO R1$$

Then, the direct oxidation of SO₂ by long lived sCI:

 $sCI + SO_2 \xrightarrow{k_2} SO_3 + RCHO$ R2

Or the alternative oxidation of SO₂ by OH radicals formed from decomposition of (stabilized) Criegee Intermediates:

80

$sCI \xrightarrow{\kappa_3} \cdot OH + R_1 COR_2$	R3
$CI \xrightarrow{k_4} \cdot OH + R_1 COR_2$	R4
$SO_2 + \cdot OH + O_2 \xrightarrow{k_5} HO_2 \cdot + SO_3$	R5

where k_i is the rate constant of each reaction, φ is the yield of sCI in the ozonolysis of alkenes, and CI is the chemically activated Criegee intermediate. The lifetime of SO₃ due to its fast reaction with H₂O to form SA is approximately 2×10^{-4} s 85 (see detailed calculation in supplement Section S1), which indicates that this reaction is so fast that almost all SO₃ will be instantaneously converted to SA. In this case, the oxidation of SO_2 is the rate-limiting step in the formation of SA. Currently, only limited types of sCI has been studied: isoprene-derived sCI (Neeb et al., 1997;Zhang et al., 2002;Atkinson et al., 2006; Newland et al., 2015b), monoterpene-derived sCI (Hatakeyama et al., 1984; Rickard et al., 1999; Zhang and Zhang, 2005; Mauldin et al., 2012; Sipila et al., 2014; Vereecken et al., 2017) and the simplest sCI including CH₂COO, CH₃CHOO and 90 (CH₃)₂COO (Hatakeyama et al., 1984;Hasson et al., 2001;Welz et al., 2012;Taatjes et al., 2013;Welz et al., 2014;Newland et al., 2015a; Vereecken et al., 2017). Based on the aforementioned studied, the yield φ of sCI can vary from 0.1 to 0.65 and the rate constants for different reactions span over several orders of magnitude, for k_1 from 1.6×10^{-18} to 2.5×10^{-16} cm³s⁻¹ and for k_2 from 1.4×10^{-13} to 2.2×10^{-10} cm³s⁻¹. The yield of OH radical from ozonolysis of different type of alkenes also covers a wide range with values of 0.68 - 0.91, 0.24 - 0.35, 0.25 - 0.44, 0.32 - 0.40 and 0.33 - 1.00 for α -pinene, β -pinene, 95 isoprene, propene and other C4-C6 alkenes respectively (Atkinson et al., 1992;Aschmann, 1993;Chew and Atkinson, 1996;Rickard et al., 1999;Siese et al., 2001;Witter et al., 2002;Berndt et al., 2003;Aschmann et al., 2003;Nguyen et al., 2009; Malkin et al., 2010). Moreover, the bimolecular reaction and decomposition reactivity of sCI is highly structuredependent. sCI with more complicated substituent groups tend to react with H₂O more slowly (Huang et al., 2015), decompose

faster (Fenske et al., 2000; Hasson et al., 2001) and more likely to react with SO2. There were also studies showing that the

100 reactions between sCI and SO₂ were pressure and temperature dependent and were commonly affected by the presence of water and other constituents (Kotzias et al., 1990;Sipila et al., 2014).

3. Ambient Observations

105

The continuous and comprehensive measurements were conducted at the west campus of Beijing University of Chemical Technology (39.95 °N', 116.31 °E'). Here we investigate the time period from 18th January to 16th March 2019. The measuring instruments are located on the fifth floor, which is about 15 m above the ground level. This station is a typical urban site, which is around 130 m to the nearest Zizhuyuan Road, 550 m to the West Third Ring Road, and surrounded by commercial properties and residential dwellings.

3.1 Measurement of Sulfuric Acid with CI-APi-TOF

Sulfuric acid is measured by a long time-of-flight chemical ionization mass specter (LTOF-CIMS, Aerodyne Research, Inc.) equipped with a nitrate chemical ionization source. The basic working principle of this instrument can be found elsewhere (Jokinen et al., 2012). In our measurement, we draw air through a stainless-steel tube with a length of 1.6 m and a diameter of 3/4 inch with a flowrate at 7.2 L min⁻¹. In addition, we have implemented a flush plate (Karsa Inc.) to effectively remove water molecules entering the instrument, which is found necessary to maintain a continuous measurement.

The quantification of sulfuric acid is derived from the ratio of bisulfate ions (with counting rates unit in ions \cdot s⁻¹) relative to 115 primary ions as follows:

$$[H_2SO_4] = \frac{HSO_4^- + H_2SO_4NO_3^-}{NO_3^- + HNO_3NO_3^- + (HNO_3)_2NO_3^-} \times C$$

The calibration factor, C, is determined from direct calibration by injecting gaseous sulfuric acid of known amounts into the instrument (Kurten et al., 2012). The diffusional wall loss of the 1.6 sampling line is 0.2423, and after taking into account of it, we a value of 6.07×10^9 cm⁻³ as the final calibration coefficient.

Six alkenes are analyzed in this study, i.e., including propylene, butylene, butadiene, isoprene, pentene and hexene, which

were detected by a single photon ionization time-of-flight mass spectrometer (SPI-MS 3000, Guangzhou Hexin Instrument

120 **3.2 Measurement of Alkenes with SPI-MS**

125

Co., Ltd., China) (Gao et al., 2013). It should be mentioned that this instrument cannot distinguish conformers, and therefore the pentene and haxene could also be cyclopentane and cyclohexane, respectively. A polydimethylsiloxane (PDMS) membrane sampling system is used. As the PDMS membrane has better selective adsorption to volatile organic compounds (VOCs), VOC molecules can be concentrated after diffusing and desorbing from the membrane under vacuum sampling condition. In this way, the detection limit of VOCs can be improved. Then the gas molecules are guided to an ionization chamber through a 2 mm-diameter stainless steel capillary, where VOC molecules are ionized by the vacuum ultraviolet (VUV) light with an ionization energy smaller than 10.8eV. For the detection of positive ions, two microchannel plates (MCPs, Hamamatsu, Japan)

130

135

assembled with a chevron-type configuration are employed. An analog to digital converter (ADC) was used to measure and record the output current signal from the MCPs.

Alkenes concentrations are quantified by performing a direct calibration. The PAMS (Photochemical Assessment Monitoring Stations) and TO-15 environmental gases (including 57 and 65 types of VOCs separately, Linde Gas North America LLC, USA) are used as two standard gases with ultra-high-purity nitrogen as the carrier gas. Gas with different concentrations of VOC standards is produced by mixing a constant carrier gas with standard gas of varying flow rates. The calibration coefficient is further calculated from the ratio between the actual concentration and the ion intensity.

3.3 Other Ancillary Measurements

The number concentration of clusters with the size range of 1.30~2.45 nm was measured with a Particle Sizer Magnifier (PSM) (Vanhanen et al., 2011), and the integrated number concentration of particles from PSM is referred as Nsub-3nm in the 140 following. The number size distributions of aerosol particles from 6 to 840 nm and from 0.52 to 19.81 µm are measured by the Differential Mobility Particle Sizer (DMPS) (Aalto et al., 2001) and the Aerodynamic Particle Sizer (APS) (Armendariz and Leith, 2002) respectively. Meteorological parameters are measured with a weather station (AWS310, Vaisala Inc.) located on the rooftop of the building. These parameters include the ambient temperature, relative humidity (RH), pressure, visibility, UVB radiation, as well as horizontal wind speed and direction. Trace gas concentrations of carbon monoxide (CO), sulfur 145 dioxide (SO₂) nitrogen oxides (NO_x) and ozone (O₃) are monitored using four Thermo Environmental Instruments (models 48i, 43i-TLE, 42i, 49i, respectively). Calibration of these instruments are performed monthly using the standard gases of known concentrations. The mass concentration of PM_{2.5} is directly measured with a Tapered Element Oscillating Microbalance Dichotomous Ambient Particulate Monitor (TEOM 1405-DF, Thermo Fisher Scientific Inc, USA) with a total flow rate of 16.67 L/min. In addition, the loss rate of gas-phase sulfuric acid described by condensation sink (CS) is calculated based on the size distribution data from DMPS and APS (Kulmala et al., 2001). 150

4. Results and Discussions

4.1 Definition of Nighttime Sulfuric Acid Event

An overview of our measurements during 18th January to 16th March 2019 is shown in Fig. 1. As our measurement period overlaps with the heating period in Beijing (from 15th November 2018 to 15th March 2019), the SO₂ level during the measurement period was higher than that of other periods in one year (from 16th March to 31st May 2019) (Fig. S1 in the Supplement).

In this work, the nighttime window is defined between 20:00 and 04:00 (following day) to exclude any possible influence of photochemistry. Fig. 2 shows the diurnal variation of SA concentration on one typical SA event night (14 March 2019) and one typical SA non-event night (3 February 2019). Overall, nighttime SA concentrations vary between 3.0×10^5 and 3.0×10^6

160 cm⁻³ in our measurement period. The nighttime event in Fig. 2 shows a distinct SA peak at around 22:00 with a maximum SA concentration of around 3.0×10⁶ cm⁻³, which is almost half of the daily maximum value. While in the non-event case, SA continues decreasing throughout the evening, reaching a minimal value of 3.0×10⁵ cm⁻³. A nighttime SA event is defined when both of the following two criteria are both met: (a) there is a distinct peak during the nighttime hours, (b) the SA concentration exceeds 1.0×10⁶ cm⁻³. The nights without distinct peaks are classified as SA non-event nights, and if a peak is identified but it does not meet criterion (b), the night is classified as an undefined night. Out of all 56 nights studied, there are in total 18 SA event nights, 16 non-event nights, and 22 undefined nights (listed in Table. S1). Thus, the overall frequency of nighttime SA events during our observation period is 32%, which means that nearly a third of nights during our observation period had distinct nighttime sulfuric acid peaks.

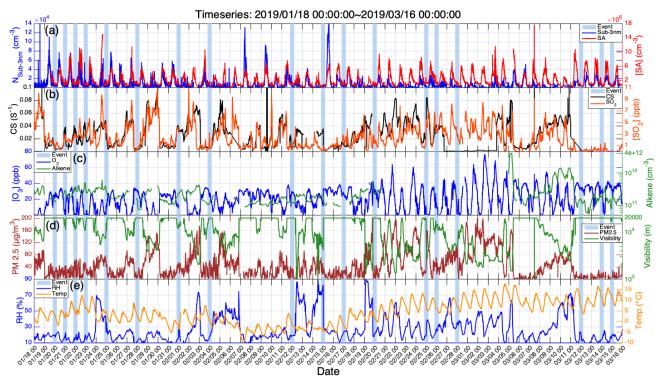


Fig. 1. Overview of different parameters measured from 18th January, 2019 to and 16th March, 2019 for (a) SA concentration and particle number concentration of sub-3nm particles (N_{Sub-3nm}, measured by PSM), (b) CS and SO₂ concentration, (c) concentration of O₃ and Alkenes, (d) PM_{2.5} and visibility, and (e) relative humidity (RH) and temperature. The light blue bars represent nights with nighttime SA events.

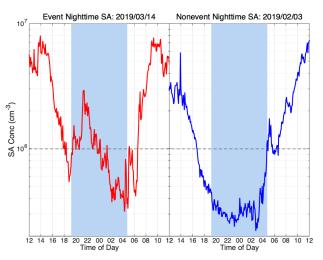


Fig. 2 Daily variation of SA concentration on a typical night with a nighttime SA event (red line, 14th March, 2019) and on a non-event night (blue line, 3rd February, 2019). The shaded blue area shows the period that is considered as nighttime in this study.

4.2 Features of Nighttime Sulfuric Acid Event

We further analyzed the features of nighttime SA event nights based on the above mentioned 18 event nights and 16 nonevent nights (Fig. 3). On SA event nights, the mass concentration of $PM_{2.5}$, the mixing ratio of NO_x , CS and RH were clearly lower, and visibility was clearly higher than on non-event nights. This suggests that nighttime SA events tend to occur under clean conditions. In addition, higher concentrations of O_3 were associated with SA event nights, whereas the concentration of SO₂ did not vary as much between SA event and non-event nights. This indicates that at most of the time, the concentration of SO₂ is not the dominant factor that explains the variation of nighttime SA.

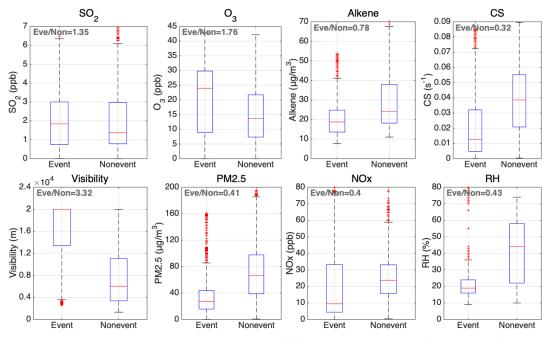
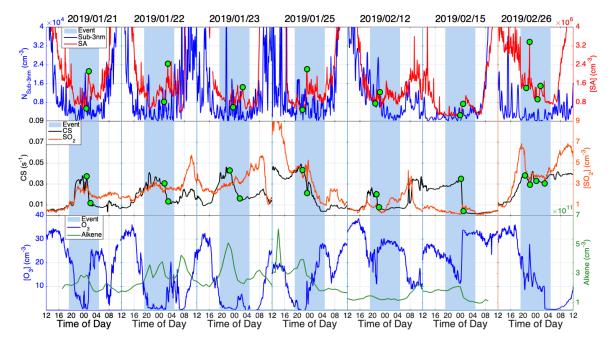


Fig. 3 Boxplots for the concentrations of SO₂, O₃ and alkenes, CS, visibility, PM_{2.5}, NO_x concentration and RH during nighttime SA event and non-event nights. In each plot, the red line is the median value, the bottom and the top blue lines are the 25 and 75 percentiles, and the whisker range covers the $\pm 2.7\sigma$ of the data. Outliers are the ones out of the $\pm 2.7\sigma$ range of all selected data. The dark gray values on the top left corners are the ratios between median values of event and non-event days.

We further investigated the determining factor for the occurrence of SA events by looking into different variables during the SA event nights. CS measurements were available for 13 event nights, during which 15 SA peaks were observed. In general, we found that eight events (53%) were mainly associated with the decrease of CS. This is demonstrated in Fig. 4, where the nighttime events as well as the simultaneous decrease of CS are highlighted with green dots. Four other cases (27%) were mostly due to the increase of SO₂ concentration (Fig. S2), and the remaining three cases were likely synergistically caused by SO₂, O₃, alkenes, CS and other parameters. More details are provided in Table S2.



195 **Fig. 4** Daily time-series of different variables on nighttime SA event days when SA events occurred under CS decrease conditions. The top panel shows the N_{Sub-3nm} and SA concentration, the middle panel shows CS and SO₂ concentration, and the bottom panel shows the concentration of O₃ and alkenes. Green dots show times when CS started to drop and reached its minimum value.

4.3 Source and Sink Balance for Nighttime Sulfuric Acid: Importance of Alkene Ozonolysis

As discussed above, nighttime SA events mainly occurred under clean conditions with low CS values. Therefore, we classified all the nighttime data set into three groups according to the air pollution level, which is assessed by visibility. The division standards for pollution level is explained in detail in Section S3. The clean (named Clean-1), mildly polluted and heavy polluted conditions are defined by visibility values which are larger than 12.0 km, in the range of 4.0 - 12.0 km and smaller than 4.0 km respectively. Accordingly, data points under each condition took up 48%, 25% and 27% of all data points.

- After classifying the data set into groups based on the pollution level, the balance between SA source and sink for each
 group was investigated separately. The ozonolysis of alkenes under dark conditions is capable of generating sCI as well as OH
 radical, both of which are able to oxidize SO₂ to form gaseous SA. However, the yields of both sCI and recycled OH radical
 remain largely unquantified. Therefore, we do not attempt to distinguish the contribution of sCI and OH radical on SA
 formation in this study, but rather treat them as a "bulk oxidant" and use an empirical parameter k_{app} to account for both
 oxidation pathways. Accordingly, the source term (production rate) of SA can be expressed as k_{app}·[SO₂]·[O₃]·[Alkene], where
 k_{app} is an overall empirical parameter that takes into account the yields of OH radical and sCI as well as the rate constants of
 their reactions with SO₂. The sink term (loss rate) consists of two parts: the condensation of SA onto particles ([SA]·CS) and
 the collision of SA monomers with each other to form SA dimers (β·[SA]²). In reality, SA monomer also collides with SA
 dimers and larger clusters, but due to the low concentration of SA clusters, those collisions are negligible compared to other
 losses. In a polluted environment where strong stabilizers of SA exist, the formation rate of stable SA dimer is close to the
- collision limit (Yao et al., 2018). Therefore, β can be taken as the hard-sphere collision rate, which is calculated to be 3.46×10^{-10} cm³ s⁻¹ (Seinfeld and Pandis, 2016). Under pseudo-steady-state (PSS) assumption (see Section S2 for detailed disscussion about PSS assumption), the source-sink balance of SA can be expressed as follows:

 $k_{app} \cdot [Alkene] \cdot [O_3] \cdot [SO_2] = [SA] \cdot CS + \beta \cdot [SA]^2$

Fig. 5 shows the nighttime correlation between SA source term and sink term under different pollution levels, with the data
 binned by SA sources. If the source and the sink term correlate, then the slope represents the overall apparent rate constant (k_{app}) concerning the reaction between oxidants (OH radical and sCI) from ozonolysis of alkenes and SO₂ for this specific pollution level.

Under Clean-1 conditions (Fig. 5 (a)), the source term and sink term have a good linear correlation (R²=0.97) in the source range of 1.0×10³³ - 4.5×10³³ (cm⁻³)³, while the balance is broken up outside of this range. These uncorrelated data points outside of this range appear when the visibility was smaller than 16.0 km (Fig. S4 (a)) along with higher concentrations of NO_x and NO ([NO_x] > ~ 40 ppb, Fig. S4 (b) and ([NO] > ~ 3 ppb, Fig. S4 (c)). High NO_x levels always relate to pollution, and NO will consume O₃, leading to much lower O₃ concentration (marked by blue empty circles in Fig. S4 (a)). Thus, we redefined the criterion for clean condition (Clean-2) so that visibility needs to be larger than 16.0 km and O₃ concentration higher than 2.0×10¹¹ cm⁻³ (~ 7 ppb). These conditions account for 38% of all data. Fig. 5 (b) shows the good correlation between [SO₂]·[O₃]·[Alkene] source term and [SA]·CS+β·[SA]² sink term (R²=0.97) under the redefined clean condition over the entire source range. This suggests that the ozonolysis of alkenes indeed have a dominant contribution to the formation of SA during nighttime under very clean conditions. Generally, the sink term of SA condensation onto particles took up 95.5% for Clean-2 condition and increase to 99.7% for heavy polluted condition. The fitted value, 95% confidence bounds, uncertainty of k_{app} and correlation coefficient R² for Clean-1and Clean-2 condition are listed in Table S3.

If we compare the SA source and sink correlation between the Clean-1 (Fig. 5 (a)) and Clean-2 (Fig. 5 (b)) condition, it is obvious that the slope of the linear region of Clean-1 condition data points (2.7×10⁻³⁰ cm⁶ s⁻²) matches well with the slope of Clean-2 condition data points (2.6×10⁻³⁰ cm⁶s⁻²), which further confirms the reliability of the balance between [SO₂]·[O₃]·[Alkene] source term and [SA]·CS+ β·[SA]² sink term under clean condition. Then we then took data from another period to further evaluate the reliability of the proposed source and sink balance. For spring period from 20th March to 20th May 2019 (Fig. S7 (b)), there is also a good linear correlation (R²=0.98) when source term is smaller than 6.0×10³³ (cm⁻³)³ with k_{app} of 1.5×10⁻³⁰ cm⁶s⁻². Although the fitted k_{app} values deviate beween these two periods, they are in the same order of magnitude. During spring period, apart from the above-mentioned six alkenes, biogenic emitted monoterpenes, which cannot be measured by our instrument and therefore are not included in the source term, start to have a bigger contribution, which likely leads to the deviation of k_{app}. Besides, the yield of sCI and the rate constant between sCI and SO₂ are to some extent temperature-dependent (Berndt et al., 2014), which may further explain at least a part of the difference of k_{app} between winter and spring observations.

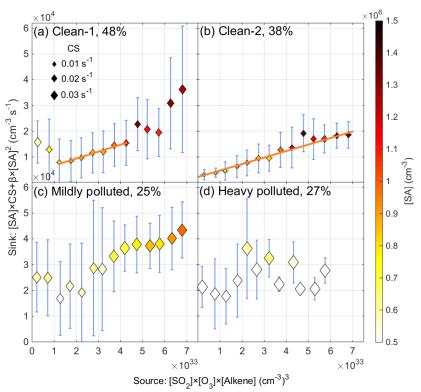


Fig. 5 Correlation between the source term ($[SO_2] \cdot [O_3] \cdot [Alkene]$) and sink term ($[SA] \cdot CS + \beta \cdot [SA]^2$) of SA under PSS assumption during nighttime (20:00-04:00) from 18th January to 16th March, 2019 for (a) Clean-1 condition, (b) Clean-2 condition, (c) mildly polluted condition and (d) heavy polluted condition. Note that the data points are mean values of corresponding bin ranges instead of the original, high time resolution data (Fig. S5). The error bars are the standard deviation of all data points in each bin.

Under mildly polluted conditions (Fig 5 (c)), the source and sink term also have a good linear correlation when source value exceeds 2.5×10³³ (cm⁻³)³, while under heavy polluted conditions, the [SO₂]·[O₃]·[Alkene] source term and [SA]·CS+β·[SA]² sink term do not show a strong correlation (Fig. 5 (d)). Most likely, this suggests that the source term cannot fully represent the actual SA source for heavy polluted conditions. For instance, there are likely additional sources of SA, such as direct emission from diesel vehicles, oil refineries, SA plants, and any other factories that use coal as heating or power supply (Srivastava et al., 2004;Arnold et al., 2006;Ahn et al., 2011;Roy et al., 2014;Sarnela et al., 2015;Godunov et al., 2017). Another possible cause for the correlation deviation under polluted condition, and to some extent also mildly polluted condition, is that the distribution of alkenes may not be constant for measurements classified into the same pollution level, that is, the k_{app} is not constant. At very clean nights when alkenes sources are considered more local and stable, dramatic changes in alkene distribution are not expected. Moreover, our instrument is only capable of measuring a limited amount of alkene species and the fitted parameter k_{app} might be overestimated.

It should be pointed out that we are not able to further deconvolute the contribution of OH radical and sCI based on the ambient observation. A well-tuned box model is a useful tool to resolve it and verify the role of the ozonolysis of alkenes on

265 the nighttime SA formation. However, such a modeling work is not included in our study, as the lacking of a complete VOC datasets in our measurement and the largely uncertain yields of sCI from the ozonolysis of various alkenes have caused challenges in ensuring the precision of the box-model.

4.4 Atmospheric Implication: Contribution of Nighttime Sulfuric Acid to Sub-3nm Particles

We show that the ozonolysis of alkenes is the major source for the considerable amount of SA that exists at night, at least 270 under unpolluted conditions. And it is found that increasing SA concentration coincided with increasing number concentration of sub-3nm particles (Fig. 6), suggesting that SA had a strong enhancement in the formation of newly formed particles, which is consistent with previous study (Cai et al., 2017). Different from SA, there was a negative correlation between the concentration of highly oxygenated organic molecules (HOMs) and N_{Sub-3nm} for both nighttime and daytime (see Fig. S8 in the supplement), indicating that HOMs were not the main driver for the formation of sub-3nm particles. Then, these elevated SA 275 concentration has the dominant contribution to the formation of sub-3nm particles in the nighttime of winter Beijing. This phenomenon is in contrast with some previous observations in forested areas where oxidation products of biogenic VOCs, especially monoterpene, were the main contributor to the formation of clusters (Eerdekens et al., 2009;Lehtipalo et al., 2011; Kammer et al., 2018; Rose et al., 2018). Both observation (Rose et al., 2018) and laboratory experiment (Lehtipalo et al., 2018) have shown that HOM dimers with extremely low volatility play a key role in the initial formation of clusters; however 280 at our site in urban Beijing, high level of NOx inhibited the production of HOM dimers and most HOMs are monomers, which have minor importance to the formation of sub-3nm particles. Nighttime sub-3nm particles has also been observed at suburban site (Kecorius et al., 2015) or areas which are strongly influenced by coastal air masses (Yu et al., 2014; Salimi et al., 2017), but the underlying mechanism are still unclear.

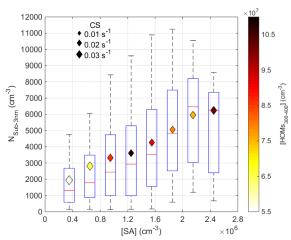


Fig. 6 Correlation between N_{Sub-3nm} and SA concentration during nighttime (20:00-04:00) from 18th January to 16th March 2019. The binned diamonds are colored by number concentration of HOMs (m/Q = 300 - 400 Th) and the size is related to CS. The red line is the median value, the bottom and the top blue lines are the 25 and 75 percentiles, and the whisker ranges cover the $\pm 2.7\sigma$ of those data in each bin.

5. Conclusions

Continuous SA measurement was conducted during the heating-supply period in urban Beijing. Frequent nighttime SA events were found and accounted for about 32 % of the total measurement nights. Most nighttime SA events were observed under unpolluted conditions and associated with a distinct drop of CS. We show that the SA source corresponding to the 11 product of O_3 , alkenes and SO_2 concentrations correlates well with the SA sink for clean conditions, and to some extent also for mildly polluted conditions. Therefore, we suggest that nighttime SA formation under these conditions can be largely attributed to the ozonolysis of alkenes which leads to the production of OH radicals as well as sCI that are able to act as

295

oxidants for SO₂. However, further deconvolution of the contribution of OH radicals, sCI and each possible alkene precursor was not possible within this study due to the inability to directly measure OH, sCI and the entire range of alkene precursors. It should also be pointed out that, under polluted conditions, there were very likely additional SA sources other than the ozonolysis of alkenes, such as direct emission from diesel vehicles, oil refineries and SA plants. Furthermore, we showed that these elevated SA had a dominant contribution to the formation of sub-3nm particles in the nighttime of winter Beijing.

300

305

Acknowledgements

This project has received funding from the National Natural Science Foundation of China (41877306), the ERC advanced grant No. 742206, the European Union's Horizon 2020 research and innovation program under grant agreement No. 654109, the Academy of Finland Center of Excellence Project No. 27204. Kaspar R. Daellenbach acknowledges the support by the SNF mobility grant P2EZP2 181599.

Data availability: Data and codes are available upon asking the corresponding author (<u>chao.yan@helsinki.fi</u>) **Author contribution:** YG, CY designed the study; YG, CY, CL, YZ, LD collected the data and materials; YG, CY, DS wrote the manuscript; all coauthors contributed to the scientific discussion and commented on the manuscript.

310 **Competing interests:** The authors declare no competing interests.

References

Aalto, P., Hämeri, K., Becker, E., Weber, R., Salm, J., Mäkelä, J. M., Hoell, C., O'dowd, C. D., Hansson, H.-C., Väkevä, M.,
Koponen, I. K., Buzorius, G., and Kulmala, M.: Physical characterization of aerosol particles during nucleation events, Tellus
B: Chemical and Physical Meteorology, 53, 344-358, 10.3402/tellusb.v53i4.17127, 2001.

- Ahn, J. Y., Okerlund, R., Fry, A., and Eddings, E. G.: Sulfur trioxide formation during oxy-coal combustion, International Journal of Greenhouse Gas Control, 5, S127-S135, 10.1016/j.ijggc.2011.05.009, 2011.
 Almeida, J., Schobesberger, S., Kuerten, A., Ortega, I. K., Kupiainen-Maatta, O., Praplan, A. P., Adamov, A., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Donahue, N. M., Downard, A., Dunne, E., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Guida, R., Hakala, J., Hansel, A., Heinritzi, M., Henschel, H., Jokinen, T., Junninen, H., Kajos,
- M., Kangasluoma, J., Keskinen, H., Kupc, A., Kurten, T., Kvashin, A. N., Laaksonen, A., Lehtipalo, K., Leiminger, M., Leppa, J., Loukonen, V., Makhmutov, V., Mathot, S., McGrath, M. J., Nieminen, T., Olenius, T., Onnela, A., Petaja, T., Riccobono, F., Riipinen, I., Rissanen, M., Rondo, L., Ruuskanen, T., Santos, F. D., Sarnela, N., Schallhart, S., Schnitzhofer, R., Seinfeld, J. H., Simon, M., Sipila, M., Stozhkov, Y., Stratmann, F., Tome, A., Troestl, J., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y.,

Virtanen, A., Vrtala, A., Wagner, P. E., Weingartner, E., Wex, H., Williamson, C., Wimmer, D., Ye, P., Yli-Juuti, T., Carslaw,

- K. S., Kulmala, M., Curtius, J., Baltensperger, U., Worsnop, D. R., Vehkamaki, H., and Kirkby, J.: Molecular understanding of sulphuric acid-amine particle nucleation in the atmosphere, Nature, 502, 359-+, 10.1038/nature12663, 2013.
 Armendariz, A. J., and Leith, D.: Concentration measurement and counting efficiency for the aerodynamic particle sizer 3320, Journal of Aerosol Science, 33, 133-148, <u>https://doi.org/10.1016/S0021-8502(01)00152-5</u>, 2002.
 Arnold, F., Pirjola, L., Aufmhoff, H., Schuck, T., Lahde, T., and Hameri, K.: First gaseous sulfuric acid measurements in
- automobile exhaust: Implications for volatile nanoparticle formation, Atmospheric Environment, 40, 7097-7105, 10.1016/j.atmosenv.2006.06.038, 2006.

Aschmann, S. M.: OH Radical Production from the Gas-Phase Reaction of O3 with a Series of Alkenes under Atmospheric Conditions, Environmental Science and Technology, 27, 1357-1363, 1993.

- Aschmann, S. M., Tuazon, E. C., Arey, J., and Atkinson, R.: Products of the Gas-Phase Reaction of O3 with Cyclohexene, The Journal of Physical Chemistry A, 107, 2247-2255, 10.1021/jp022122e, 2003.
- Atkinson, R., Aschmann, S. M., Arey, J., and Shorees, B.: Formation of OH radicals in the gas phase reactions of O3 with a series of terpenes, Journal of Geophysical Research: Atmospheres, 97, 6065-6073, 10.1029/92JD00062, 1992.

335

- Atkinson, R., Baulch, D. L., Cox, R. A., Crowley, J. N., Hampson, R. F., Hynes, R. G., Jenkin, M. E., Rossi, M. J., and Troe,
 J.: Evaluated kinetic and photochemical data for atmospheric chemistry: Volume II gas phase reactions of organic species,
 Atmospheric Chemistry and Physics, 6, 3625-4055, 10,5194/acp-6-3625-2006, 2006.
- Berndt, T., Böge, O., and Stratmann, F.: Gas-phase ozonolysis of α-pinene: gaseous products and particle formation, Atmospheric Environment, 37, 3933-3945, <u>https://doi.org/10.1016/S1352-2310(03)00501-6</u>, 2003.
 Berndt, T., Jokinen, T., Sipila, M., Mauldin, R. L., III, Herrmann, H., Stratmann, F., Junninen, H., and Kulmala, M.: H2SO4 formation from the gas-phase reaction of stabilized Criegee Intermediates with SO2: Influence of water vapour content and
- temperature, Atmospheric Environment, 89, 603-612, 10.1016/j.atmosenv.2014.02.062, 2014.
 Birmili, W., Berresheim, H., Plass-Dulmer, C., Elste, T., Gilge, S., Wiedensohler, A., and Uhrner, U.: The Hohenpeissenberg aerosol formation experiment (HAFEX): a long-term study including size-resolved aerosol, H2SO4, OH, and monoterpenes measurements, Atmospheric Chemistry and Physics, 3, 361-376, 10.5194/acp-3-361-2003, 2003.
 Boy, M., Karl, T., Turnipseed, A., Mauldin, R. L., Kosciuch, E., Greenberg, J., Rathbone, J., Smith, J., Held, A., Barsanti, K.,
- Wehner, B., Bauer, S., Wiedensohler, A., Bonn, B., Kulmala, M., and Guenther, A.: New particle formation in the front range of the colorado rocky mountains, Atmospheric Chemistry and Physics, 8, 1577-1590, 10.5194/acp-8-1577-2008, 2008.
 Cai, R., Yang, D., Fu, Y., Wang, X., Li, X., Ma, Y., Hao, J., Zheng, J., and Jiang, J.: Aerosol surface area concentration: a governing factor in new particle formation in Beijing, Atmospheric Chemistry and Physics, 17, 12327-12340, 10.5194/acp-17-12327-2017, 2017.
- Chew, A. A., and Atkinson, R.: OH radical formation yields from the gas-phase reactions of O3 with alkenes and monoterpenes, Journal of Geophysical Research: Atmospheres, 101, 28649-28653, 10.1029/96JD02722, 1996.
 Dada, L., Ylivinkka, I., Baalbaki, R., Li, C., Guo, Y., Yan, C., Yao, L., Sarnela, N., Jokinen, T., Daellenbach, K. R., Yin, R., Deng, C., Chu, B., Nieminen, T., Kontkanen, J., Stolzenburg, D., Sipilä, M., Hussein, T., Paasonen, P., Bianchi, F., Salma, I., 13

Weidinger, T., Pikridas, M., Sciare, J., Jiang, J., Liu, Y., Petäjä, T., Kerminen, V. M., and Kulmala, M.: Sources and sinks

driving sulphuric acid concentrations in contrasting environments: implications on proxy calculations, Atmos. Chem. Phys.
 Discuss., 2020, 1-24, 10.5194/acp-2020-155, 2020.

365

2017.

Eerdekens, G., Yassaa, N., Sinha, V., Aalto, P. P., Aufmhoff, H., Arnold, F., Fiedler, V., Kulmala, M., and Williams, J.: VOC measurements within a boreal forest during spring 2005: on the occurrence of elevated monoterpene concentrations during night time intense particle concentration events, Atmospheric Chemistry and Physics, 9, 8331-8350, 10.5194/acp-9-8331-2009, 2009.

Eisele, F. L., Lovejoy, E. R., Kosciuch, E., Moore, K. F., Mauldin, R. L., Smith, J. N., McMurry, P. H., and Iida, K.: Negative atmospheric ions and their potential role in ion-induced nucleation, Journal of Geophysical Research-Atmospheres, 111, 10.1029/2005jd006568, 2006.

Erupe, M. E., Benson, D. R., Li, J., Young, L.-H., Verheggen, B., Al-Refai, M., Tahboub, O., Cunningham, V., Frimpong, F.,

- Viggiano, A. A., and Lee, S.-H.: Correlation of aerosol nucleation rate with sulfuric acid and ammonia in Kent, Ohio: An atmospheric observation, Journal of Geophysical Research-Atmospheres, 115, 10.1029/2010jd013942, 2010.
 Fenske, J. D., Hasson, A. S., Ho, A. W., and Paulson, S. E.: Measurement of Absolute Unimolecular and Bimolecular Rate Constants for CH3CHOO Generated by the trans-2-Butene Reaction with Ozone in the Gas Phase, The Journal of Physical Chemistry A, 104, 9921-9932, 10.1021/jp0016636, 2000.
- 375 Fiedler, V., Dal Maso, M., Boy, M., Aufmhoff, H., Hoffmann, J., Schuck, T., Birmili, W., Hanke, M., Uecker, J., Arnold, F., and Kulmala, M.: The contribution of sulphuric acid to atmospheric particle formation and growth: a comparison between boundary layers in Northern and Central Europe, Atmospheric Chemistry and Physics, 5, 1773-1785, 10.5194/acp-5-1773-2005, 2005.
- Gao, W., Tan, G., Hong, Y., Li, M., Nian, H., Guo, C., Huang, Z., Fu, Z., Dong, J., Xu, X., Cheng, P., and Zhou, Z.:
 Development of portable single photon ionization time-of-flight mass spectrometer combined with membrane inlet, International Journal of Mass Spectrometry, 334, 8-12, 10.1016/j.ijms.2012.09.003, 2013.
 Godunov, E. B., Kuznetsova, I. A., and Klevleev, V. M.: Comprehensive Utilization of Spent Sulfuric Acid from Petrochemical Industries and Chemical Current Sources, Chemical and Petroleum Engineering, 53, 531-533, 10.1007/s10556-017-0376-9,
- Hasson, A. S., Orzechowska, G., and Paulson, S. E.: Production of stabilized Criegee intermediates and peroxides in the gas phase ozonolysis of alkenes 1. Ethene, trans-2-butene, and 2,3-dimethyl-2-butene, Journal of Geophysical Research-Atmospheres, 106, 34131-34142, 10.1029/2001jd000597, 2001.
 Hatakeyama, S., Kobayashi, H., and Akimoto, H.: Gas-phase oxidation of sulfur dioxide in the ozone-olefin reactions, The Journal of Physical Chemistry, 88, 4736-4739, 10.1021/j150664a058, 1984.
- 390 Huang, H.-L., Chao, W., and Lin, J. J.-M.: Kinetics of a Criegee intermediate that would survive high humidity and may oxidize atmospheric SO2, Proceedings of the National Academy of Sciences of the United States of America, 112, 10857-10862, 10.1073/pnas.1513149112, 2015.

Iida, K., Stolzenburg, M. R., McMurry, P. H., and Smith, J. N.: Estimating nanoparticle growth rates from size-dependent 14

charged fractions: Analysis of new particle formation events in Mexico City, Journal of Geophysical Research-Atmospheres,

395 113, 10.1029/2007jd009260, 2008.

Jokinen, T., Sipila, M., Junninen, H., Ehn, M., Lonn, G., Hakala, J., Petaja, T., Mauldin, R. L., III, Kulmala, M., and Worsnop, D. R.: Atmospheric sulphuric acid and neutral cluster measurements using CI-APi-TOF, Atmospheric Chemistry and Physics, 12, 4117-4125, 10.5194/acp-12-4117-2012, 2012.

Jokinen, T., Sipila, M., Kontkanen, J., Vakkari, V., Tisler, P., Duplissy, E. M., Junninen, H., Kangasluoma, J., Manninen, H. E.,

- Petaja, T., Kulmala, M., Worsnop, D. R., Kirkby, J., Virkkula, A., and Kerminen, V. M.: Ion-induced sulfuric acid-ammonia nucleation drives particle formation in coastal Antarctica, Science Advances, 4, 10.1126/sciadv.aat9744, 2018.
 Junninen, H., Hulkkonen, M., Riipinen, I., Nieminen, T., Hirsikko, A., Suni, T., Boy, M., Lee, S.-H., Vana, M., Tammet, H., Kerminen, V.-M., and Kulmala, M.: Observations on nocturnal growth of atmospheric clusters, Tellus Series B-Chemical and Physical Meteorology, 60, 365-371, 10.1111/j.1600-0889.2008.00356.x, 2008.
- 405 Kammer, J., Perraudin, E., Flaud, P. M., Lamaud, E., Bonnefond, J. M., and Villenave, E.: Observation of nighttime new particle formation over the French Landes forest, Science of the Total Environment, 621, 1084-1092, 10.1016/j.scitotenv.2017.10.118, 2018.

Kecorius, S., Zhang, S., Wang, Z., Groess, J., Ma, N., Wu, Z., Ran, L., Hu, M., Wang, P., Ulevicius, V., and Wiedensohler, A.: NOCTURNAL AEROSOL PARTICLE FORMATION IN THE NORTH CHINA PLAIN, Lithuanian Journal of Physics, 55,

410 44-53, 2015.

Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagne, S., Ickes, L., Kuerten, A., Kupc,
A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S., Tsagkogeorgas, G., Wimmer, D., Amorim, A., Bianchi, F.,
Breitenlechner, M., David, A., Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S., Hansel, A., Hauser, D., Jud, W.,
Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R., Makhmutov, V., Mathot, S.,

- Mikkila, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A., Pereira, P., Petaja, T., Schnitzhofer, R., Seinfeld, J. H., Sipila, M., Stozhkov, Y., Stratmann, F., Tome, A., Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E., Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger, U., and Kulmala, M.: Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol nucleation, Nature, 476, 429-U477, 10.1038/nature10343, 2011.
 Kotzias, D., Fytianos, K., and Geiss, F.: Reaction of monoterpenes with ozone, sulphur dioxide and nitrogen dioxide—gas-
- 420 phase oxidation of SO2 and formation of sulphuric acid, Atmospheric Environment. Part A. General Topics, 24, 2127-2132, https://doi.org/10.1016/0960-1686(90)90246-J, 1990.
 - Kuerten, A., Jokinen, T., Simon, M., Sipila, M., Sarnela, N., Junninen, H., Adamov, A., Almeida, J., Amorim, A., Bianchi, F.,
 Breitenlechner, M., Dommen, J., Donahue, N. M., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Hakala, J., Hansel, A.,
 Heinritzi, M., Hutterli, M., Kangasluoma, J., Kirkby, J., Laaksonen, A., Lehtipalo, K., Leiminger, M., Makhmutov, V., Mathot,
- S., Onnela, A., Petaja, T., Praplan, A. P., Riccobono, F., Rissanen, M. P., Rondo, L., Schobesberger, S., Seinfeld, J. H., Steiner, G., Tome, A., Troestl, J., Winkler, P. M., Williamson, C., Wimmer, D., Ye, P., Baltensperger, U., Carslaw, K. S., Kulmala, M., Worsnop, D. R., and Curtius, J.: Neutral molecular cluster formation of sulfuric acid-dimethylamine observed in real time under atmospheric conditions, Proceedings of the National Academy of Sciences of the United States of America, 111, 15019-15

Kuerten, A., Bergen, A., Heinritzi, M., Leiminger, M., Lorenz, V., Piel, F., Simon, M., Sitals, R., Wagner, A. C., and Curtius, J.: Observation of new particle formation and measurement of sulfuric acid, ammonia, amines and highly oxidized organic molecules at a rural site in central Germany, Atmospheric Chemistry and Physics, 16, 12793-12813, 10.5194/acp-16-12793-2016, 2016.

Kulmala, M., Maso, M. D., Mäkelä, J. M., Pirjola, L., Väkevä, M., Aalto, P., Miikkulainen, P., Hämeri, K., and O'dowd, C.

D.: On the formation, growth and composition of nucleation mode particles, Tellus B: Chemical and Physical Meteorology, 53, 479-490, 10.3402/tellusb.v53i4.16622, 2001.
Kulmala, M., Lehtinen, K. E. J., and Laaksonen, A.: Cluster activation theory as an explanation of the linear dependence between formation rate of 3nm particles and sulphuric acid concentration, Atmospheric Chemistry and Physics, 6, 787-793, 10.5194/acp-6-787-2006, 2006.

- Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T., Petäjä, T., Sipilä, M., Schobesberger, S., Rantala, P., Franchin, A., Jokinen, T., Järvinen, E., Äijälä, M., Kangasluoma, J., Hakala, J., Aalto, P. P., Paasonen, P., Mikkilä, J., Vanhanen, J., Aalto, J., Hakola, H., Makkonen, U., Ruuskanen, T., Mauldin, R. L., Duplissy, J., Vehkamäki, H., Bäck, J., Kortelainen, A., Riipinen, I., Kurtén, T., Johnston, M. V., Smith, J. N., Ehn, M., Mentel, T. F., Lehtinen, K. E. J., Laaksonen, A., Kerminen, V.-M., and Worsnop, D. R.: Direct Observations of Atmospheric Aerosol Nucleation, Science, 339,
- 445 943, 10.1126/science.1227385, 2013.

Kurten, A., Rondo, L., Ehrhart, S., and Curtius, J.: Calibration of a Chemical Ionization Mass Spectrometer for the Measurement of Gaseous Sulfuric Acid, Journal of Physical Chemistry A, 116, 6375-6386, 10.1021/jp212123n, 2012.

Lee, S. H., Gordon, H., Yu, H., Lehtipalo, K., Haley, R., Li, Y. X., and Zhang, R. Y.: New Particle Formation in the Atmosphere: From Molecular Clusters to Global Climate, Journal of Geophysical Research-Atmospheres, 124, 7098-7146, 10.1029/2018jd029356, 2019.

Lehtipalo, K., Sipila, M., Junninen, H., Ehn, M., Berndt, T., Kajos, M. K., Worsnop, D. R., Petaja, T., and Kulmala, M.: Observations of Nano-CN in the Nocturnal Boreal Forest, Aerosol Science and Technology, 45, 499-509, 10.1080/02786826.2010.547537, 2011.

Lehtipalo, K., Yan, C., Dada, L., Bianchi, F., Xiao, M., Wagner, R., Stolzenburg, D., Ahonen, L. R., Amorim, A., Baccarini,

- A., Bauer, P. S., Baumgartner, B., Bergen, A., Bernhammer, A.-K., Breitenlechner, M., Brilke, S., Buchholz, A., Mazon, S. B.,
 Chen, D., Chen, X., Dias, A., Dommen, J., Draper, D. C., Duplissy, J., Ehn, M., Finkenzeller, H., Fischer, L., Frege, C., Fuchs,
 C., Garmash, O., Gordon, H., Hakala, J., He, X., Heikkinen, L., Heinritzi, M., Helm, J. C., Hofbauer, V., Hoyle, C. R., Jokinen,
 T., Kangasluoma, J., Kerminen, V.-M., Kim, C., Kirkby, J., Kontkanen, J., Kuerten, A., Lawler, M. J., Mai, H., Mathot, S.,
 Mauldin, R. L., III, Molteni, U., Nichman, L., Nie, W., Nieminen, T., Ojdanic, A., Onnela, A., Passananti, M., Petaja, T., Piel,
- F., Pospisilova, V., Quelever, L. L. J., Rissanen, M. P., Rose, C., Sarnela, N., Schallhart, S., Schuchmann, S., Sengupta, K., Simon, M., Sipila, M., Tauber, C., Tome, A., Trostl, J., Vaisanen, O., Vogel, A. L., Volkamer, R., Wagner, A. C., Wang, M., Weitz, L., Wimmer, D., Ye, P., Ylisirnio, A., Zha, Q., Carslaw, K. S., Curtius, J., Donahue, N. M., Flagan, R. C., Hansel, A., Riipinen, I., Virtanen, A., Winkler, P. M., Baltensperger, U., Kulmala, M., and Worsnop, D. R.: Multicomponent new particle 16

formation from sulfuric acid, ammonia, and biogenic vapors, Science Advances, 4, 10.1126/sciadv.aau5363, 2018.

465 Lu, Y., Yan, C., Fu, Y., Chen, Y., Liu, Y., Yang, G., Wang, Y., Bianchi, F., Chu, B., Zhou, Y., Yin, R., Baalbaki, R., Garmash, O., Deng, C., Wang, W., Liu, Y., Petaja, T., Kerminen, V.-M., Jiang, J., Kulmala, M., and Wang, L.: A proxy for atmospheric daytime gaseous sulfuric acid concentration in urban Beijing, Atmospheric Chemistry and Physics, 19, 1971-1983, 10.5194/acp-19-1971-2019, 2019.

Malkin, T. L., Goddard, A., Heard, D. E., and Seakins, P. W.: Measurements of OH and HO2 yields from the gas phase 470 ozonolysis of isoprene, Atmospheric Chemistry and Physics, 10, 1441-1459, 10.5194/acp-10-1441-2010, 2010.

Mauldin, R. L., Eisele, F. L., Tanner, D. J., Kosciuch, E., Shetter, R., Lefer, B., Hall, S. R., Nowak, J. B., Buhr, M., Chen, G., Wang, P., and Davis, D.: Measurements of OH, H2SO4, and MSA at the South Pole during ISCAT, Geophysical Research Letters, 28, 3629-3632, 10.1029/2000gl012711, 2001.

Mauldin, R. L., III, Berndt, T., Sipilae, M., Paasonen, P., Petaja, T., Kim, S., Kurten, T., Stratmann, F., Kerminen, V. M., and

475 Kulmala, M.: A new atmospherically relevant oxidant of sulphur dioxide, Nature, 488, 193-196, 10.1038/nature11278, 2012. Mazon, S. B., Kontkanen, J., Manninen, H. E., Nieminen, T., Kerminen, V.-M., and Kulmala, M.: A long-term comparison of nighttime cluster events and daytime ion formation in a boreal forest, Boreal Environment Research, 21, 242-261, 2016. McMurry, P. H., Fink, M., Sakurai, H., Stolzenburg, M. R., Mauldin, R. L., Smith, J., Eisele, F., Moore, K., Sjostedt, S., Tanner, D., Huey, L. G., Nowak, J. B., Edgerton, E., and Voisin, D.: A criterion for new particle formation in the sulfur-rich Atlanta 480

atmosphere, Journal of Geophysical Research-Atmospheres, 110, 10.1029/2005jd005901, 2005.

- Myllys, N., Chee, S., Olenius, T., Lawler, M., and Smith, J.: Molecular-Level Understanding of Synergistic Effects in Sulfuric Acid-Amine-Ammonia Mixed Clusters, Journal of Physical Chemistry A, 123, 2420-2425, 10.1021/acs.jpca.9b00909, 2019. Neeb, P., Sauer, F., Horie, O., and Moortgat, G. K.: Formation of hydroxymethyl hydroperoxide and formic acid in alkene ozonolysis in the presence of water vapour, Atmospheric Environment, 31, 1417-1423, https://doi.org/10.1016/S1352-
- 485 2310(96)00322-6, 1997.

Newland, M. J., Rickard, A. R., Alam, M. S., Vereecken, L., Munoz, A., Rodenas, M., and Bloss, W. J.: Kinetics of stabilised Criegee intermediates derived from alkene ozonolysis: reactions with SO2, H2O and decomposition under boundary layer conditions, Physical Chemistry Chemical Physics, 17, 4076-4088, 10.1039/c4cp04186k, 2015a.

Newland, M. J., Rickard, A. R., Vereecken, L., Munoz, A., Rodenas, M., and Bloss, W. J.: Atmospheric isoprene ozonolysis: impacts of stabilised Criegee intermediate reactions with SO2, H2O and dimethyl sulfide, Atmospheric Chemistry and Physics, 490 15, 9521-9536, 10.5194/acp-15-9521-2015, 2015b.

Nguyen, T. L., Peeters, J., and Vereecken, L.: Theoretical study of the gas-phase ozonolysis of beta-pinene (C10H16), Physical Chemistry Chemical Physics, 11, 5643-5656, 10.1039/b822984h, 2009.

Paasonen, P., Sihto, S.-L., Nieminen, T., Vuollekoski, H., Riipinen, I., Plass-Dulmer, C., Berresheim, H., Birmili, W., and 495 Kulmala, M.: Connection between new particle formation and sulphuric acid at Hohenpeissenberg (Germany) including the influence of organic compounds, Boreal Environment Research, 14, 616-629, 2009. Paasonen, P., Nieminen, T., Asmi, E., Manninen, H. E., Petaja, T., Plass-Duelmer, C., Flentje, H., Birmili, W., Wiedensohler,

A., Horrak, U., Metzger, A., Hamed, A., Laaksonen, A., Facchini, M. C., Kerminen, V. M., and Kulmala, M.: On the roles of 17

sulphuric acid and low-volatility organic vapours in the initial steps of atmospheric new particle formation, Atmospheric

- Chemistry and Physics, 10, 11223-11242, 10.5194/acp-10-11223-2010, 2010.
 Riccobono, F., Schobesberger, S., Scott, C. E., Dommen, J., Ortega, I. K., Rondo, L., Almeida, J., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Downard, A., Dunne, E. M., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Hansel, A., Junninen, H., Kajos, M., Keskinen, H., Kupc, A., Kuerten, A., Kvashin, A. N., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Petaja, T., Praplan, A. P., Santos, F. D., Schallhart, S., Seinfeld, J. H., Sipila, M.,
- Spracklen, D. V., Stozhkov, Y., Stratmann, F., Tome, A., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y., Vrtala, A., Wagner,
 P. E., Weingartner, E., Wex, H., Wimmer, D., Carslaw, K. S., Curtius, J., Donahue, N. M., Kirkby, J., Kulmala, M., Worsnop,
 D. R., and Baltensperger, U.: Oxidation Products of Biogenic Emissions Contribute to Nucleation of Atmospheric Particles,
 Science, 344, 717-721, 10.1126/science.1243527, 2014.
 - Rickard, A. R., Johnson, D., McGill, C. D., and Marston, G.: OH Yields in the Gas-Phase Reactions of Ozone with Alkenes, The Journal of Physical Chemistry A, 103, 7656-7664, 10.1021/jp9916992, 1999.
- Riipinen, I., Sihto, S. L., Kulmala, M., Arnold, F., Dal Maso, M., Birmili, W., Saarnio, K., Teinila, K., Kerminen, V. M., Laaksonen, A., and Lehtinen, K. E. J.: Connections between atmospheric sulphuric acid and new particle formation during QUEST III-IV campaigns in Heidelberg and Hyyti, Atmospheric Chemistry and Physics, 7, 1899-1914, 10.5194/acp-7-1899-2007, 2007.
- 515 Rose, C., Zha, Q., Dada, L., Yan, C., Lehtipalo, K., Junninen, H., Mazon, S. B., Jokinen, T., Sarnela, N., Sipila, M., Petaja, T., Kerminen, V.-M., Bianchi, F., and Kulmala, M.: Observations of biogenic ion-induced cluster formation in the atmosphere, Science Advances, 4, 10.1126/sciadv.aar5218, 2018.

Roy, B., Chen, L., and Bhattacharya, S.: Nitrogen Oxides, Sulfur Trioxide, and Mercury Emissions during Oxyfuel Fluidized Bed Combustion of Victorian Brown Coal, Environmental Science & Technology, 48, 14844-14850, 10.1021/es504667r, 2014.

- Salimi, F., Rahman, M. M., Clifford, S., Ristovski, Z., and Morawska, L.: Nocturnal new particle formation events in urban environments, Atmos. Chem. Phys., 17, 521-530, 10.5194/acp-17-521-2017, 2017.
 Sarnela, N., Jokinen, T., Nieminen, T., Lehtipalo, K., Junninen, H., Kangasluoma, J., Hakala, J., Taipale, R., Schobesberger, S., Sipila, M., Larnimaa, K., Westerholm, H., Heijari, J., Kerminen, V.-M., Petaja, T., and Kulmala, M.: Sulphuric acid and aerosol particle production in the vicinity of an oil refinery, Atmospheric Environment, 119, 156-166,
- 525 10.1016/j.atmosenv.2015.08.033, 2015.

510

530

Seinfeld, J. H., and Pandis, S. N.: Atmospheric chemistry and physics: from air pollution to climate change, John Wiley & Sons, 2016.

Siese, M., Becker, K. H., Brockmann, K. J., Geiger, H., Hofzumahaus, A., Holland, F., Mihelcic, D., and Wirtz, K.: Direct Measurement of OH Radicals from Ozonolysis of Selected Alkenes: A EUPHORE Simulation Chamber Study, Environmental Science & Technology, 35, 4660-4667, 10.1021/es010150p, 2001.

Sipila, M., Jokinen, T., Berndt, T., Richters, S., Makkonen, R., Donahue, N. M., Mauldin, R. L., III, Kurten, T., Paasonen, P., Sarnela, N., Ehn, M., Junninen, H., Rissanen, M. P., Thornton, J., Stratmann, F., Herrmann, H., Worsnop, D. R., Kulmala, M., Kerminen, V. M., and Petaja, T.: Reactivity of stabilized Criegee intermediates (sCIs) from isoprene and monoterpene 18 ozonolysis toward SO2 and organic acids, Atmospheric Chemistry and Physics, 14, 12143-12153, 10.5194/acp-14-12143-

535 2014, 2014.

555

Srivastava, R. K., Miller, C. A., Erickson, C., and Jambhekar, R.: Emissions of sulfur trioxide from coal-fired power plants, Journal of the Air & Waste Management Association, 54, 750-762, 10.1080/10473289.2004.10470943, 2004.
Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M. M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M.: Climate Change 2013: The physical science basis. contribution of working group I to the fifth assessment report of IPCC

- the intergovernmental panel on climate change, in, Cambridge University Press, 2014.
 Taatjes, C. A., Welz, O., Eskola, A. J., Savee, J. D., Scheer, A. M., Shallcross, D. E., Rotavera, B., Lee, E. P. F., Dyke, J. M., Mok, D. K. W., Osborn, D. L., and Percival, C. J.: Direct Measurements of Conformer-Dependent Reactivity of the Criegee Intermediate CH3CHOO, Science, 340, 177-180, 10.1126/science.1234689, 2013.
 Taipale, R., Sarnela, N., Rissanen, M., Junninen, H., Rantala, P., Korhonen, F., Siivola, E., Berndt, T., Kulmala, M., Mauldin,
- R. L., III, Petaja, T., and Sipila, M.: New instrument for measuring atmospheric concentrations of non-OH oxidants of SO2, Boreal Environment Research, 19, 55-70, 2014.
 Vanhanen, J., Mikkila, J., Lehtipalo, K., Sipila, M., Manninen, H. E., Siivola, E., Petaja, T., and Kulmala, M.: Particle Size

Vereecken, L., Novelli, A., and Taraborrelli, D.: Unimolecular decay strongly limits the atmospheric impact of Criegee intermediates, Physical Chemistry Chemical Physics, 19, 31599-31612, 10.1039/c7cp05541b, 2017.

Magnifier for Nano-CN Detection, Aerosol Science and Technology, 45, 533-542, 10.1080/02786826.2010.547889, 2011.

Wang, Z. B., Hu, M., Yue, D. L., Zheng, J., Zhang, R. Y., Wiedensohler, A., Wu, Z. J., Nieminen, T., and Boy, M.: Evaluation on the role of sulfuric acid in the mechanisms of new particle formation for Beijing case, Atmospheric Chemistry and Physics, 11, 12663-12671, 10.5194/acp-11-12663-2011, 2011.

Weber, R. J., Marti, J. J., McMurry, P. H., Eisele, F. L., Tanner, D. J., and Jefferson, A.: Measurements of new particle formation and ultrafine particle growth rates at a clean continental site, 102, 4375-4385, 10.1029/96jd03656, 1997.

- Weber, R. J., McMurry, P. H., Mauldin, L., Tanner, D. J., Eisele, F. L., Brechtel, F. J., Kreidenweis, S. M., Kok, G. L., Schillawski, R. D., and Baumgardner, D.: A study of new particle formation and growth involving biogenic and trace gas species measured during ACE 1, 103, 16385-16396, 10.1029/97jd02465, 1998.
- Weber, R. J., McMurry, P. H., Mauldin III, R. L., Tanner, D. J., Eisele, F. L., Clarke, A. D., and Kapustin, V. N.: New Particle
 Formation in the Remote Troposphere: A Comparison of Observations at Various Sites, 26, 307-310, 10.1029/1998gl900308, 1999.

Welz, O., Savee, J. D., Osborn, D. L., Vasu, S. S., Percival, C. J., Shallcross, D. E., and Taatjes, C. A.: Direct Kinetic Measurements of Criegee Intermediate (CH2OO) Formed by Reaction of CH2I with O-2, Science, 335, 204-207, 10.1126/science.1213229, 2012.

Welz, O., Eskola, A. J., Sheps, L., Rotavera, B., Savee, J. D., Scheer, A. M., Osborn, D. L., Lowe, D., Booth, A. M., Xiao, P., Khan, M. A. H., Percival, C. J., Shallcross, D. E., and Taatjes, C. A.: Rate Coefficients of C1 and C2 Criegee Intermediate Reactions with Formic and Acetic Acid Near the Collision Limit: Direct Kinetics Measurements and Atmospheric Implications, Angewandte Chemie-International Edition, 53, 4547-4550, 10.1002/anie.201400964, 2014.

Witter, M., Berndt, T., Böge, O., Stratmann, F., and Heintzenberg, J.: Gas-phase ozonolysis: Rate coefficients for a series of

- 570 terpenes and rate coefficients and OH yields for 2-methyl-2-butene and 2,3-dimethyl-2-butene, International Journal of Chemical Kinetics, 34, 394-403, 10.1002/kin.10063, 2002.
 - Yao, L., Garmash, O., Bianchi, F., Zheng, J., Yan, C., Kontkanen, J., Junninen, H., Mazon, S. B., Ehn, M., Paasonen, P., Sipila,
 M., Wang, M. Y., Wang, X. K., Xiao, S., Chen, H. F., Lu, Y. Q., Zhang, B. W., Wang, D. F., Fu, Q. Y., Geng, F. H., Li, L.,
 Wang, H. L., Qiao, L. P., Yang, X., Chen, J. M., Kerminen, V. M., Petaja, T., Worsnop, D. R., Kulmala, M., and Wang, L.:
- 575 Atmospheric new particle formation from sulfuric acid and amines in a Chinese megacity, Science, 361, 278-+, 10.1126/science.aao4839, 2018.

Yu, H., McGraw, R., and Lee, S.-H.: Effects of amines on formation of sub-3 nm particles and their subsequent growth, Geophysical Research Letters, 39, 10.1029/2011gl050099, 2012.

Yu, H., Hallar, A. G., You, Y., Sedlacek, A., Springston, S., Kanawade, V. P., Lee, Y. N., Wang, J., Kuang, C. G., McGraw, R.

- L., McCubbin, I., Mikkila, J., and Lee, S. H.: Sub-3nm particles observed at the coastal and continental sites in the United States, Journal of Geophysical Research-Atmospheres, 119, 860-879, 10.1002/2013jd020841, 2014.
 Zhang, D., Lei, W. F., and Zhang, R. Y.: Mechanism of OH formation from ozonolysis of isoprene: kinetics and product yields, Chemical Physics Letters, 358, 171-179, 10.1016/s0009-2614(02)00260-9, 2002.
 Zhang, D., and Zhang, R.: Ozonolysis of alpha-pinene and beta-pinene: Kinetics and mechanism, Journal of Chemical Physics,
- 585 122, 10.1063/1.1862616, 2005.