## 1 Characterization of the light absorbing properties, chromophores composition

- 2 and sources of brown carbon aerosol in Xi'an, Northwest China
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# 25 Abstract

26 The impact of brown carbon aerosol (BrC) on the Earth's radiative forcing balance has

27 been widely recognized but remains uncertain, mainly because the relationships among BrC 28 sources, chromophores, and optical properties of aerosol are poorly understood. In this work, 29 the light absorption properties and chromophore composition of BrC were investigated for 30 samples collected in Xi'an, Northwest China from 2015 to 2016. Both absorption Ångström exponent and mass absorption efficiency show distinct seasonal differences, which could be 31 attributed to the differences in sources and chromophore composition of BrC. Three groups of 32 light-absorbing organics were found to be important BrC chromophores, including those show 33 34 multiple absorption peaks at wavelength > 350 nm (12 polycyclic aromatic hydrocarbons and their derivatives) and those show single absorption peak at wavelength < 350 nm (10 35 nitrophenols and nitrosalicylic acids and 3 methoxyphenols). These measured BrC 36 chromophores show distinct seasonal differences and contribute on average about 1.1% and 3.3% 37 of light absorption of methanol-soluble BrC at 365 nm in summer and winter, respectively, 38 39 about 7 and 5 times higher than the corresponding carbon mass fractions in total organic carbon. 40 The sources of BrC were resolved by positive matrix factorization (PMF) using these chromophores instead of commonly used non-light absorbing organic markers as model inputs. 41 42 Our results show that in spring vehicular emissions and secondary formation are major sources of BrC (~70%), in fall coal combustion and vehicular emissions are major sources (~70%), in 43 winter biomass burning and coal combustion become major sources (~80%), while in summer 44 45 secondary BrC dominates (~60%).

## 46 **1 Introduction**

Brown carbon (BrC) is an important component of atmospheric aerosol particles and has 47 significant effects on radiative forcing and climate (Feng et al., 2013; Laskin et al., 2015; Zhang 48 49 et al., 2017a). BrC can efficiently absorb solar radiation and reduce the photolysis rates of atmospheric radicals (Jacobsan, 1999; Li et al., 2011; Mok et al., 2016), which ultimately 50 influences the atmospheric photochemistry process, the formation of secondary organic aerosol 51 (SOA), and therefore the regional air quality (Mohr et al., 2013; Laskin et al., 2015; Moise et 52 53 al., 2015). In addition, some components in BrC, such as nitrated aromatic compounds (NACs) 54 (Teich et al., 2017; Wang et al., 2018) and polycyclic aromatic hydrocarbons (PAHs) (Samburova et al., 2016; Huang et al., 2018), have adverse effects on human health (Bandowe 55

et al., 2014; Shen et al., 2018). The significant effects of BrC on environment, climate, air
quality and living things call for more studies to understand its chemical characteristics, sources
and the links with optical properties.

Investigating the chemical composition of BrC at molecular level is necessary, because 59 60 even small amounts of compounds can have a significant effect on the light absorption 61 properties of BrC and profound atmospheric implication (Mohr et al., 2013; Zhang et al., 2013; Teich et al., 2017; Huang et al., 2018). A number of studies have investigated the BrC 62 63 composition at molecular level (Mohr et al., 2013; Zhang et al., 2013; Chow et al., 2015; Samburova et al., 2016; Lin et al., 2016, 2017, 2018; Teich et al., 2017; Huang et al., 2018; Lu 64 65 et al., 2019). For example, Zhang et al. (2013) measured 8 NACs in Los Angeles and found that they contributed about 4% of water-soluble BrC absorption at 365 nm. Huang et al. (2018) 66 measured 18 PAHs and their derivatives in Xi'an and found that they accounted for on average 67  $\sim 1.7\%$  of the overall absorption of methanol-soluble BrC. A state-of-the-art high performance 68 69 liquid chromatography-photodiode array-high resolution mass spectrometry (HPLC-PDA-70 HRMS) was applied to investigate the elemental composition of BrC chromophores in biomass 71 burning aerosol (Lin et al., 2016, 2017, 2018). Despite these efforts, the molecular composition 72 of atmospheric BrC still remains largely unknown due to its complexity in emission sources 73 and formation processes.

74 Field observations and laboratory studies show that BrC has various sources, including 75 primary emissions such as combustion and secondary formation from various atmospheric 76 processes (Laskin et al., 2015). Biomass burning, including forest fires and burning of crop 77 residues, is considered as the main source of BrC (Teich et al., 2017; Lin et al., 2017). Coal 78 burning and vehicle emissions are also important primary sources of BrC (Yan et al., 2017; Xie 79 et al., 2017; Sun et al., 2017; Li et al., 2019; Song et al., 2019). Secondary BrC is produced 80 through multiple-phase reactions occurring in or between gas phase, particle phase, and cloud droplets. For example, nitrification of aromatic compounds (Harrison et al., 2005; Lu et al., 81 2011), oligomers of acid-catalyzed condensation of hydroxyl aldehyde (De Haan et al., 2009; 82 83 Shapiro et al., 2009), and reaction of ammonia (NH<sub>3</sub>) or amino acids with carbonyls (De Haan 84 et al., 2011; Nguyen et al., 2013; Flores et al., 2014) can all produce BrC. Condensed phase reactions and aqueous-phase reactions have also been found to be important formation pathways for secondary BrC in ambient air (Gilardoni et al., 2016). In addition, atmospheric aging processes can lead to either enhancement or bleaching of the BrC absorption (Lambe et al., 2013; Lee et al., 2014; Zhong and Jang, 2014), further challenging the characterization of BrC.

90 As the starting point of the Silk Road, Xi'an is an important inland city in northwestern China experiencing severe particulate air pollution, especially during heating period with 91 92 enhanced coal combustion and biomass burning activities (Wang et al., 2016; Ni et al., 2018). In this study, we performed spectroscopic measurement and chemical analysis of PM<sub>2.5</sub> filter 93 samples in Xi'an to investigate: 1) seasonal variations in the light absorption properties and 94 95 chromophore composition of BrC, and their relationships; 2) sources of BrC in different seasons based on positive matrix factorization (PMF) model with light-absorbing organic markers as 96 97 input species.

## 98 2 Experimental

#### 99 2.1 Aerosol sampling

A total of 112 daily ambient PM2.5 filter samples were collected on pre-baked (780 °C, 3 100 h) quartz-fiber filters (20.3 × 25.4 cm, Whatman, QM-A, Clifton, NJ, USA) in November-101 102 December 2015, April-May, July, October-November 2016, representing winter, spring, 103 summer and fall, respectively. Filter samples were collected using a Hi-Vol PM<sub>2.5</sub> air sampler (Tisch, Cleveland, OH) at a flow rate of 1.05 m<sup>3</sup> min<sup>-1</sup> on the roof (~10 m above ground level, 104 105 34.22°N, 109.01°E) of the Institute of Earth Environment, Chinese Academy of Sciences, 106 which was surrounded by residential areas without large industrial activities. After collection, 107 the filter samples were wrapped in baked aluminum foils and stored in a freezer (-20 °C) until 108 further analysis.

# 109 2.2 Light absorption measurement

110 One punch of loaded filter (0.526 cm<sup>2</sup>) was taken from each sample and sonicated for 30 111 minutes in 10 mL of ultrapure water (> 18.2 M $\Omega$  · cm) or methanol (HPLC grade, J. T. Baker, 112 Phillipsburg, NJ, USA). The extracts were then filtered with a 0.45 µm PTFE pore syringe filter to remove insoluble materials. The light absorption spectra of water-soluble and methanolsoluble BrC were measured with an UV-Vis spectrophotometer (300-700 nm) equipped with a liquid waveguide capillary cell (LWCC-3100, World Precision Instrument, Sarasota, FL, USA) following the method by Hecobian et al. (2010). The measured absorption data can be converted to the absorption coefficient  $Abs_{\lambda}$  (M m<sup>-1</sup>) by equation (1):

118 
$$Abs_{\lambda} = (A_{\lambda} - A_{700}) \frac{V_{l}}{V_{a} \times L} \times \ln(10)$$
 (1)

119 where  $A_{700}$  is the absorption at 700 nm, serving as a reference to account for baseline drift,  $V_1$ 120 is the volume of water or methanol that the filter was extracted into,  $V_a$  is the volume of sampled 121 air, L is the optical path length (0.94 m). A factor of ln(10) is used to convert the log base-10 122 (recorded by UV-Vis spectrophotometer) to natural logarithm to provide base-e absorption 123 coefficient. The absorption coefficient of water-soluble or methanol-soluble organics at 365 nm 124 (Abs<sub>365</sub>) is used to represent water-soluble or methanol-soluble BrC absorption, respectively.

125 The mass absorption efficiency (MAE:  $m^2 gC^{-1}$ ) of BrC in the extracts can be calculated 126 as:

127 
$$MAE_{\lambda} = \frac{Abs_{\lambda}}{M}$$
(2)

where M ( $\mu$ gC m<sup>-3</sup>) is the concentration of water-soluble organic carbon (WSOC) for water extracts or methanol-soluble organic carbon (MSOC) for methanol extracts. Note that organic carbon (OC) is often used to replace MSOC because direct measurement of MSOC is technically difficult and many studies have shown that most of OC (~ 90%) can be extracted by methanol (Chen and Bond, 2010; Cheng et al., 2016; Xie et al., 2019).

133 The wavelength-dependent light absorption of chromophores in solution, termed as
134 absorption Ångström exponent (AAE), can be described as:

 $Abs_{\lambda} = K \cdot \lambda^{-AAE}$ (3)

136 where K is a constant related to the concentration of chromophores and AAE is calculated by 137 linear regression of log Abs<sub> $\lambda$ </sub> versus log  $\lambda$  in the wavelength range of 300-410 nm.

#### 138 **2.3 Chemical analysis**

OC was measured with a thermal/optical carbon analyzer (DRI, model 2001) following
the IMPROVE-A protocol (Chow et al., 2011). WSOC was measured with a TOC/TN analyzer

141 (TOC-L, Shimadzu, Japan) (Ho et al., 2015).

Organic compounds listed in Table S1 were analyzed with a gas chromatograph-mass 142 143 spectrometer (GC-MS, Agilent Technologies, Santa Clara, CA, USA). Prior to the GC-MS analysis, the silvlation derivatization was conducted using a routine method (e.g., Wang et al., 144 2016; Al-Naiema and stone, 2017). Briefly, a quarter of 47 mm filter sample was ultrasonically 145 146 extracted with 2 mL of methanol for 15 minutes and repeated three times. The extracts were filtered with a 0.45 µm PTFE syringe filter and then evaporated with a rotary evaporator to ~1 147 148 mL and dried with a gentle stream of nitrogen. Then, 50 µL of N,O-149 bis(trimethylsilyl)trifluoroacetamide (BSTFA-TMCS; Fluka Analytical 99%) and 10 µL of pyridine were added. The mixture was heated for 3 h at 70 °C for silvlation. After reaction, 140 150 µL of n-hexane were added to dilute the derivatives. Finally, 2 µL aliquot of the derivatized 151 extracts were introduced into the GC-MS, which was equipped with a DB-5MS column 152 153 (Agilent Technologies, Santa Clara, CA, USA), electron impact (EI) ionization source (70 eV), 154 and a GC inlet of 280 °C. The GC oven temperature was held at 50 °C for 2 min, ramped to 120 °C at a rate of 15 °C min<sup>-1</sup>, and finally reached 300 °C at a rate of 5 °C min<sup>-1</sup> (held for 16 min). 155 156 Note that the derivatization for NACs was conducted at 70 °C for 3 h which is slightly different from the protocol used in Al-Naiema and stone (2017), because symmetrical peak shapes and 157 high intensities for NACs can also be obtained under this condition in our study (see Fig. S1). 158 159 In our study, 4-nitrophenol-2,3,5,6-d4 was used as an internal standard to correct for potential 160 loss for NACs quantification (Chow et al., 2015). For the quantification of other organic 161 compounds, an external standard method was used through daily calibration with working 162 standard solutions. Also, for every 10 samples, a procedural blank and a spiked sample (i.e., 163 ambient sample spiked with known amounts of standards) were measured to check the 164 interferences and recoveries. The measured recoveries are 80-102% and the uncertainties 165 (RSDs) are < 10% for measured organic compounds.

166 **2.4 Source apportionment of BrC** 

167 Source apportionment of methanol-soluble BrC was performed using positive matrix 168 factorization (PMF) as implemented by the multilinear engine (ME-2; Paatero, 1997) via the 169 Source Finder (SoFi) interface written in Igor Wavemetrics (Canonaco et al., 2013). Abs<sub>365,MSOC</sub> 170 and those light-absorbing species including fluoranthene (FLU), pyrene (PYR), chrysene 171 (CHR), benzo(a)anthracene (BaA), benzo(a)pyrene (BaP), benzo(b)fluoranthene (BbF), 172 benzo(k)fluoranthene (BkF), indeno[1,2,3-cd]pyrene (IcdP), benzo(ghi)perylene (BghiP), 9,10anthracenequinone (9,10AQ), benzanthrone (BEN), benzo[b]fluoren-11-one (BbF11O), 173 vanillic acid, vanillin and syringyl acetone were used as model inputs, together with some 174 commonly used markers, i.e., phthalic acid, hopanes (17a(H),21B(H)-30-norhopane, 175 176  $17\alpha(H), 21\beta(H)$ -hopane,  $17\alpha(H), 21\beta(H)-(22S)$ -homohopane,  $17\alpha(H), 21\beta(H)-(22R)$ -177 homohopane, referred to as HP1-HP4, respectively), picene, and levoglucosan. The input data include species concentrations and uncertainties. The method detection limits (MDLs), 178 179 calculated as three times of the standard deviation of the blank filters, were used to estimate 180 species-specific uncertainties, following Liu et al. (2017). Furthermore, for a clear separation 181 of sources profiles, the contribution of corresponding markers was set to 0 in the sources 182 unrelated to the markers (see Table S2). This source apportionment protocol is very similar to our previous study (Huang et al., 2014). 183

184 **3 Results and discussion** 

## 185 **3.1 Light absorption properties of water- and methanol-soluble BrC**

Fig. 1 shows the temporal profiles of Abs<sub>365</sub> of water- and methanol-soluble BrC, together 186 187 with the concentrations of WSOC and OC (representing MSOC). They all show similar seasonal variations with the highest average in winter, followed by fall, spring and summer (see 188 Table S3). WSOC contributed annually  $54.4 \pm 16.2\%$  of the OC mass, with the highest 189 190 contribution in summer (66.1  $\pm$  15.5%) and the lowest contribution in winter (45.1  $\pm$  10.2%). 191 The higher WSOC fraction in OC during summer is largely contributed by SOA and to some 192 extent by biomass burning emissions because both SOA and biomass burning OA consist of 193 high fraction of WSOC (Ram et al., 2012; Yan et al., 2015; Daellenbach et al., 2016). The lower 194 WSOC fractions in OC during winter could be attributed to enhanced emissions from coal combustion which produce a large fraction of water-insoluble organics (Daellenbach et al., 195 196 2016; Yan et al., 2017). Abs<sub>365,MSOC</sub> is approximately 2 times (range 1.7-2.3) higher than 197 Abs365,WSOC, which is similar to the results measured in Beijing (Cheng et al., 2016),

southeastern Tibetan Plateau (Zhu et al., 2018), Gwangju, Korea (Park et al., 2018) and the 198 199 Research Triangle Park, USA (Xie et al., 2019), indicating that the optical properties of BrC 200 could be largely underestimated when using water as the extracting solvent as non-polar 201 fraction of BrC is also important to light absorption of BrC (Sengupta et al., 2018). In Fig. S2 202 we summarized those previously reported Abs<sub>365,WSOC</sub> (as Abs<sub>365,MSOC</sub> was not commonly 203 measured in many previous studies) values at different sites in Asian urban and remote areas 204 and the US. Abs<sub>365,WSOC</sub> is significantly higher in most Asian urban regions than in the Asian 205 remote sites and the US, and show clear seasonal variations. The high light absorption of BrC 206 in Asian urban regions, especially during winter, may have important effects on regional climate and radiation forcing (Park et al., 2010; Laskin et al., 2015). As discussed in Feng et al. (2013), 207 the average global climate forcing of BrC was estimated to be 0.04-0.11 W m<sup>-2</sup> and above 0.25 208 W m<sup>-2</sup> in urban sites of south and east Asia regions, which is about 25% of the radiative forcing 209 of black carbon (BC, 1.07 W m<sup>-2</sup>). Thus, to further understand the influence of BrC on regional 210 211 radiation forcing, it is essential to identify and quantify the sources of BrC in Asia.

212 The seasonal averages of AAE of water-soluble BrC were between 5.32 and 6.15 without 213 clear seasonal trend (see Table S3). The seasonal averages of AAE of methanol-soluble BrC 214 were relatively lower than those of water-soluble BrC, ranging from 4.45 to 5.18 which is similar to the results in Los Angeles Basin (Zhang et al., 2013) and Gwangju, Korea (Park et 215 216 al., 2018). This is because methanol can extract more compounds with high conjugation degree 217 and strong light-absorbing capability (e.g., PAHs) at longer wavelength (> 350 nm) (Samburova 218 et al., 2016). The AAE values of water-soluble BrC (as AAE of methanol-soluble BrC was not 219 commonly measured in many previous studies) in urban, rural and remote regions show a large 220 difference (see Fig. 2a), typically with much lower AAE values in urban regions than those in 221 rural and remote regions, indicating the difference in sources and chemical composition of 222 chromophores. The urban regions are mainly affected by anthropogenic emissions. Therefore, urban BrC may contain a large amount of aromatic chromophores with high conjugation degree, 223 224 which absorb light at a longer wavelength and have lower AAE values (Lambe et al., 2013; 225 Wang et al., 2018).

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The average MAE<sub>365</sub> values of water- and methanol-soluble BrC show large seasonal

variations, with highest values in winter (1.85 and 1.50 m<sup>2</sup> gC<sup>-1</sup>, respectively), followed by fall 227 (1.18 and 1.52 m<sup>2</sup> gC<sup>-1</sup>), spring (1.01 and 0.79 m<sup>2</sup> gC<sup>-1</sup>), and summer (0.91 and 1.21 m<sup>2</sup> gC<sup>-1</sup>). 228 229 Such large seasonal differences indicate seasonal difference in BrC sources. For example, 230 contributions from coal burning and biomass burning were much larger in winter than in other seasons due to large residential heating activities (also see Section 3.3 for more details). 231 232 Compared to previous studies (Fig. 2b), the average values of MAE<sub>365,WSOC</sub> are obviously higher in urban sites than in rural and remote sites that are less influenced by anthropogenic activities. 233 234 The higher MAE<sub>365,WSOC</sub> values in urban regions is likely associated with enhanced 235 anthropogenic emissions from e.g., coal combustion and biomass burning, and the lower 236 MAE<sub>365,WSOC</sub> values in rural and remote regions could be attributed to biogenic sources or aged 237 secondary BrC (Lei et al., 2018; Xie et al., 2019).

## 238 **3.2** Chemical characterization of the BrC chromophores

239 Given the complexity in emission sources and formation processes, the molecular composition of atmospheric BrC remains largely unknown. PAHs, NACs and MOPs have 240 241 recently been found as major chromophores in biomass burning-derived BrC (Lin et al., 2016, 242 2017, 2018). However, these compounds can also be directly emitted by coal combustion and motor vehicle or formed by secondary reactions (Harrison et al., 2005; Iinuma et al., 2010; Liu 243 et al., 2017; Wang et al., 2018; Lu et al., 2019), making source attribution of atmospheric BrC 244 245 more challenging. To obtain the exact molecular composition of BrC chromophores and 246 understand the influence of a specific chromophore on BrC optical property, we measured the light absorption characteristics of available chromophore standards including 12 PAHs, 10 247 NACs and 3 MOPs, and quantified their concentrations in PM2.5 samples with GC-MS. The 248 249 light absorption contribution of individual chromophores to that of methanol-soluble BrC in the wavelength range of 300-500 nm was estimated according to its concentration and mass 250 251 absorption efficiency (see Supplementary). Fig. 3 shows the contribution of carbon content in identified BrC chromophores to the total OC mass. They all show obvious seasonal variations 252 253 with the highest values in winter and lowest in summer. The seasonal difference can be up to a 254 factor of 5-6. The contribution of PAHs ranged from 0.12% in summer to 0.47% in winter, NACs from 0.02% in summer to 0.13% in winter, and MOPs from 0.01% in summer to 0.06% 255

in winter. It should be noted that NACs are dominated by 4-nitrophenol and 4-nitrocatechol in
spring, fall and winter, but by 4-nitrophenol and 5-nitrosalicylic acid in summer. The difference
is likely due to enhanced summertime formation of 5-nitrosalicylic acid, which is more oxidized
than other nitrated phenols measured in this study (Wang et al., 2018).

The seasonally averaged contributions of PAHs, NACs, MOPs and total measured 260 261 chromophores to light absorption of methanol-soluble BrC between 300 to 500 nm are shown in Fig. 4. They show large seasonal variations and wavelength dependence. Specifically, PAHs 262 263 made the largest contribution to BrC light absorption in fall, followed by winter, spring and 264 summer, and show two large absorption peaks at about 365 nm and 380 nm, which are mainly associated with the absorption of BaP, BghiP, IcdP, FLU, BkF and BaA (see Fig. S3). Compared 265 266 to PAHs, NACs show the largest contribution in winter, followed by fall, spring and summer, and exhibit only one absorption peak at about 320 nm in spring and summer and at about 330 267 268 nm in fall and winter. The red shift in the absorption peak could be attributed to the increase of 269 contributions from 4-nitrocatechol, 4-methyl-5nitrocatechol and 3-methyl-5-nitrocatechol which have absorption peak at about 330-350 nm (see Fig. S3). Different from PAHs and NACs, 270 271 MOPs contribute the most in winter, followed by spring, fall and summer, and only show one 272 absorption peak at about 310 nm. The difference in light absorption contributions of different 273 chromophores in different seasons reflects the difference in sources, emission strength and 274 atmospheric formation processes.

275 The total contributions of PAHs, NACs and MOPs to the light absorption of methanolsoluble BrC ranged from 0.47% (summer) to 1.56% (winter) at the wavelength of 300-500 nm 276 277 and ranged from 1.05% (summer) to 3.26% (winter) at the wavelength of 365 nm (see Table 1). 278 The average contribution of PAHs to the BrC light absorption at 365 nm was 0.97% in summer 279 (the lowest) and 2.69% in fall (the highest), the contribution of NACs was 0.09% in summer 280 and 0.82% in winter, and the contribution of MOPs was 0.006% in summer and 0.024% in winter. The low contributions of these measured chromophores to the light absorption of 281 methanol-soluble BrC are consistent with previous studies. For example, Huang et al. (2018) 282 283 measured 18 PAHs and their derivatives, which on average contributed  $\sim 1.7\%$  of the overall 284 absorption of methanol-soluble BrC in Xi'an. Mohr et al. (2013) estimated the contribution of

five NACs to particulate BrC light absorption at 370 nm to be ~4% in Detling, UK. Zhang et 285 286 al. (2013) measured eight NACs, which accounted for ~4% of water-soluble BrC absorption at 287 365 nm in Los Angeles. Teich et al. (2017) determined eight NACs during six campaigns at five locations in summer and winter, and founded that the mean contribution of NACs to water-288 soluble BrC absorption at 370 nm ranged from 0.10% to 1.25% under acidic conditions and 289 from 0.13% to 3.71% under alkaline conditions. Slightly different from these previous studies, 290 we investigated the contributions of three groups of chromophores with different light-291 292 absorbing properties to the light absorption of BrC, and provided further understanding in the 293 relationships between optical properties and chemical composition of BrC in the atmosphere. 294 For example, vanillin, which has negligible contribution to BrC light absorption at 365 nm, can 295 produce secondary BrC through oxidation and thus enhance the light absorption by a factor of 296 5-7 (Li et al., 2014; Smith et al., 2016). The contribution of PAHs to the light absorption of methanol-soluble BrC at 365 nm was 5-13 times that of their mass fraction of carbon in OC, 6-297 298 9 times for NACs, and 0.4-0.7 times for MOPs (4-8 times at 310 nm for MOPs). These results 299 further demonstrate that even a small amount of chromophores can have a disproportionately 300 high impact on the light absorption properties of BrC, and that the light absorption of BrC is 301 likely determined by a number of chromophores with strong light absorption ability (Kampf et al., 2012; Teich et al., 2017). Of note, a large fraction of BrC chromophores are still not 302 303 identified so far, and more studies are therefore necessary to better understand the BrC 304 chemistry. Based on laboratory and ambient studies, imidazoles (Kampf et al., 2012; Teich et 305 al., 2016), quinones (Lee et al., 2014; Pillar et al., 2017), nitrogenous PAHs (Lin et al., 2016; 306 Lin et al., 2018), polyphenols (Lin et al., 2016; Pillar et al., 2017) and oligomers with higher 307 conjugation (Lin et al., 2014; Lavi et al., 2017) could be included in future studies.

308 3.3 Sources of BrC

Two approaches have been used to quantify the sources of BrC, including multiple linear regression and receptor models such as PMF. For example, Washenfelder et al. (2015) utilized multiple linear regression to determine the contribution of individual OA factors resolved by PMF to OA light absorption in the southeastern America. Moschos et al. (2018) combined the time series of PMF-resolved OA factors with the time series of light absorption of water-soluble 314 OA extract as model inputs to quantify the sources of BrC in Magadino and Zurich, Switzerland. Xie et al. (2019) quantified the sources of BrC in southeastern America using Abs<sub>365</sub>, elemental 315 316 carbon (EC), OC, WSOC, isoprene sulfate ester, monoterpene sulfate ester, levoglucosan and 317 isoprene SOA tracers as PMF model inputs. However, it should be noted that previous studies 318 mainly rely on the correlation between measured light absorption and organic tracers that do 319 not contain a BrC chromophore, and therefore may lead to bias in BrC source apportionment. To better constrain the sources of BrC (i.e., contribution to Abs<sub>365,MSOC</sub>), we used BrC 320 321 chromophores as PMF model inputs. The inputs include vanillic acid, vanillin, and syringyl 322 acetone for BrC from biomass burning, and FLU, PYR, CHR, BaA, BaP, BbF, BkF, IcdP, BghiP, 323 for BrC from incomplete combustion and other light absorbing chromophores 9,10AQ, BEN, 324 and BbF11O. In addition, we included commonly used markers levoglucosan for biomass burning, phthalic acid for secondary BrC, hopanes for vehicle emission and picene for coal 325 326 burning in the model inputs.

Four factors were resolved, including vehicle emission, coal burning, biomass burning and 327 secondary formation. The uncertainties for PMF analysis are < 10% for secondary formation 328 329 and biomass burning, < 15% for vehicle emission and coal burning. The profile of each factor 330 is shown in Fig. S4. The first factor is characterized by a high contribution of phthalic acid, a tracer of secondary formation of OA. The second factor is dominated by hopanes, mainly from 331 332 vehicular emissions. The third factor is characterized by high contributions of PI, BaP, BbF, 333 BkF, IcdP, BghiP, mainly from coal combustion emissions, while the fourth factor has high 334 contributions of levoglucosan, vanillic acid, vanillin, syringyl acetone from biomass burning 335 emissions. The seasonal difference in relative contribution of each factor to BrC light absorption 336 is shown in Fig. 5. In spring, vehicular emissions (34%) and secondary formation (37%) were 337 the main contributors to BrC and coal combustion also had a relatively large contribution (29%). 338 In summer, secondary formation constituted the largest fraction ( $\sim 60\%$ ), mainly due to enhanced photochemical formation of secondary BrC. In fall, vehicular emissions (38%), coal 339 combustion (29%) and biomass burning (22%) all had significant contributions to BrC. In 340 341 winter, coal combustion (44%) and biomass burning (36%) were the main contributors, due to emissions from residential biomass burning (wood and crop residues) and coal combustion for 342

heating. In terms of absolute contributions to absorption of MSOC at 365 nm (see Table S4), 343 secondary formation contributed 1.75, 2.55, 1.70, 6.20 M m<sup>-1</sup> in spring, summer, fall and winter, 344 345 respectively. The high contribution in winter can be attributed to abundant precursors (volatile organic compounds) co-emitted with other primary sources (especially coal burning and 346 biomass burning), while the high contribution in summer might be due to strong photochemical 347 348 activity. For spring and fall, the absolute contributions from secondary formation were very similar, indicating moderate precursor emission and moderate photochemical activity. Also it 349 350 should be noted that the absolute contributions of vehicle emission to absorption of MSOC at 351 365 nm were still higher in spring and fall than those in summer and winter, yet these differences by a factor of 2-9 are still less pronounced than the differences (spring/fall vs. winter) for other 352 primary emissions (> 40 times for coal burning and > 25 times for biomass burning). In 353 354 particular, the high vehicle contribution in fall might be affected by high relative humidity in fall (83% in fall vs. 61-69% in other seasons, on average) resulting in high vehicular PM<sub>2.5</sub> 355 356 emissions (Chio et al., 2010). Such large seasonal difference in emission sources and atmospheric processes of BrC indicates that more studies are required to better understand the 357 358 relationship between chemical composition, formation processes, and light absorption 359 properties of BrC.

## 360 4 Conclusion

361 The light absorption properties of water- and methanol-soluble BrC in different seasons were investigated in Xi'an. The light absorption coefficient of methanol-soluble BrC was 362 approximately 2 times higher than that of water-soluble BrC at 365 nm, and had an average 363 MAE<sub>365</sub> value of  $1.27 \pm 0.46$  m<sup>2</sup> gC<sup>-1</sup>. The average MAE<sub>365</sub> value of water-soluble BrC was 1.19 364  $\pm 0.51 \text{ m}^2 \text{ gC}^{-1}$ , which is comparable to those in previous studies at urban sites but higher than 365 366 those in rural and remote areas. The seasonally averaged AAE values of water-soluble BrC ranged from 5.32 to 6.15, which are higher than those of methanol-soluble BrC (between 4.45 367 and 5.18). In combination with previous studies, we found that AAE values of water-soluble 368 369 BrC were much lower in urban regions than those in rural and remote regions. The difference of optical properties of BrC in different regions could be attributed to the difference in sources 370 371 and chemical composition of BrC chromophores. The contributions of 12 PAHs, 10 NACs and 372 3 MOPs to the light absorption of methanol-soluble BrC were determined and showed large 373 seasonal variations. Specifically, the total contribution to methanol-soluble BrC light absorption at 365 nm ranged from 1.1% to 3.3%, which is 5-7 times higher than their carbon mass fractions 374 in total OC. This result indicates that the light absorption of BrC is likely determined by an 375 amount of chromophores with strong light absorption ability. Four major sources of methanol-376 377 soluble BrC were identified, including secondary formation, vehicle emission, coal combustion and biomass burning. On average, secondary formation and vehicular emission were the main 378 379 contributors of BrC in spring (~70%). Vehicular emission (38%), coal burning (29%) and 380 biomass burning (22%) all contributed significantly to BrC in fall. Coal combustion and biomass burning were the major contributors in winter ( $\sim$ 80%), and secondary formation was 381 the predominant source in summer (~60%). The large variations of BrC sources in different 382 seasons suggest that more studies are needed to understand the seasonal difference in chemical 383 composition, formation processes, and light absorption properties of BrC, as well as their 384 385 relationships.

386 **5** Abbreviations of organics

BaA

#### PAHs (Polycyclic Aromatic Hydrocarbons) 387

- Benzo(a)anthracene 388 389 BaP Benzo(a)pyrene
- 390 **B**bF Benzo(b)fluoranthene
- BbF11O 391 Benzo[b]fluoren-11-One
- BEN 392 Benzanthrone
- 393 **BghiP** Benzo(ghi)perylene
- 394 BkF Benzo(k)fluoranthene
- 395 CHR Chrysene
- FLU Fluoranthene 396
- Indeno[1,2,3-cd]pyrene 397 IcdP
- 398 PYR Pyrene
- 399 9,10AQ 9,10-Anthracenequinone

# 400 NACs (Nitrated Aromatic Compounds)

401	2M4NP	2-Methyl-4-Nitrophenol
402	2,6DM4NP	2,6-Dimethyl-4-Nitropheol
403	3M4NP	3-Methyl-4-Nitrophenol
404	3M5NC	3-Methyl-5-Nitrocatechol
405	3NSA	3-Nitrosalicylic Acid
406	4M5NC	4-Methyl-5-Nitrocatechol
407	4NC	4-Nitrocatechol
408	4NP	4-Nitrophenol
409	4N1N	4-Nitro-1-Naphthol
410	5NSA	5-Nitrosalicylic Acid
411	MOP (Methoxyphene	ols)
411 412	<b>MOP (Methoxyphen</b> SyA	ols) Syringyl Acetone
412	SyA	Syringyl Acetone
412 413	SyA VaA	Syringyl Acetone Vanillic Acid
412 413 414	SyA VaA VAN	Syringyl Acetone Vanillic Acid
412 413 414 415	SyA VaA VAN <b>Hopanes</b>	Syringyl Acetone Vanillic Acid Vanillin
412 413 414 415 416	SyA VaA VAN <b>Hopanes</b> HP1	Syringyl Acetone Vanillic Acid Vanillin 17α(H),21β(H)-30-Norhopane

423 Supplement. The Supplement related to this article is available online at

424 *Author contributions*. RJH designed the study. Data analysis was done by WY, LY, and RJH.

*Data availability*. Raw data used in this study are archived at the Institute of Earth Environment,
Chinese Academy of Sciences, and are available on request by contacting the corresponding
author.

WY, LY and RJH interpreted data, prepared the display items and wrote the manuscript. All
authors commented on and discussed the manuscript.

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720 Table 1. Annual and seasonal mean contributions of measured PAHs, NACs and MOPs to

methanol-soluble BrC light absorption at 365 nm. Hyphens denote the measured value of more

than one third of the samples is below the dete	ection limit.
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Compounds	MAE <sub>365</sub>	Contribution to BrC light absorption at 365 nm (%)				
	(m <sup>2</sup> gC <sup>-1</sup> )	Annual	Spring	Summer	Fall	Winter
Fluoranthene (FLU)	4.25	0.11	0.05	0.02	0.05	0.15
Pyrene (PYR)	0.46	0.01	0.00	0.00	0.01	0.01
Chrysene (CHR)	0.00	0.00	0.00	0.00	0.00	0.00
Benzo(a)anthracene (BaA)	2.06	0.04	0.01	0.01	0.02	0.05
Benzo(a)pyrene (BaP)	9.31	1.04	0.76	0.39	1.16	1.10
Benzo(b)fluoranthene (BbF)	4.10	0.17	0.14	0.07	0.17	0.18
Benzo(k)fluoranthene (BkF)	3.47	0.04	0.03	0.02	0.04	0.04
Indeno[1,2,3-cd]pyrene (IcdP)	4.68	0.51	0.50	0.24	0.71	0.46
Benzo(ghi)perylene (BghiP)	8.95	0.29	0.28	0.16	0.41	0.26
9,10-Anthracenequinone (9,10AQ)	0.28	0.01	0.00	0.00	0.00	0.01
Benzanthrone (BEN)	6.13	0.11	0.08	0.05	0.11	0.12
Benzo[b]fluoren-11-one (BbF11O)	1.89	0.02	0.02	0.01	0.02	0.03
4-Nitrophenol (4NP)	2.17	0.08	0.06	0.02	0.05	0.10
4-Nitro-1-naphthol (4N1N)	9.71	-	-	-	-	0.03
2-Methyl-4-nitrophenol (2M4NP)	2.81	0.03	0.01	0.01	0.01	0.04
3-Methyl-4-nitrophenol (3M4NP)	2.65	0.02	0.01	0.00	0.01	0.03
2,6-Dimethyl-4-nitrophenol (2,6DM4NP)	3.27	-	-	-	-	0.01
4-Nitrocatechol (4NC)	7.91	0.27	0.05	0.03	0.20	0.35
3-Methyl-5-nitrocatechol (3M5NC)	5.77	-	-	-	0.05	0.11
4-Methyl-5-nitrocatechol (4M5NC)	7.29	-	-	-	0.06	0.13
3-Nitrosalicylicacid (3NSA)	3.86	-	-	-	-	0.01
5-Nitrosalicylicacid (5NSA)	3.36	0.03	0.01	0.02	0.04	0.02
Syringyl acetone (SyA)	0.25	0.01	0.01	0.00	0.01	0.01
Vanillin (VAN)	8.17	0.01	0.00	0.00	0.00	0.01
Vanillic acid (VaA)	0.66	0.00	0.00	0.00	0.00	0.00
Total	103.46	2.80	2.02	1.05	3.13	3.26

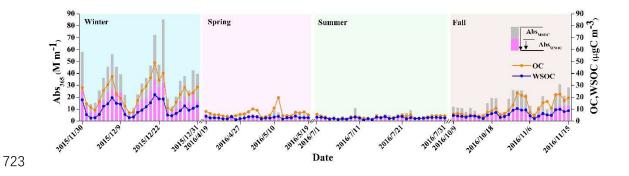


Figure 1. Time series of the light absorption coefficient of water-soluble and methanol-soluble

BrC at 365 nm (Abs<sub>365,WSOC</sub> and Abs<sub>365, MSOC</sub>, respectively), as well as OC and WSOC
concentrations.

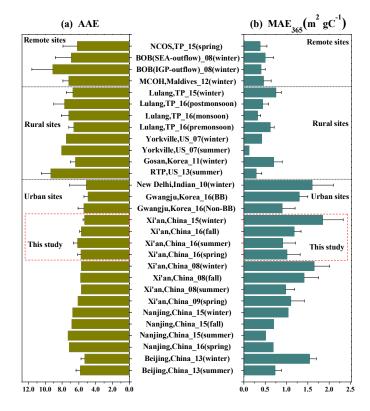
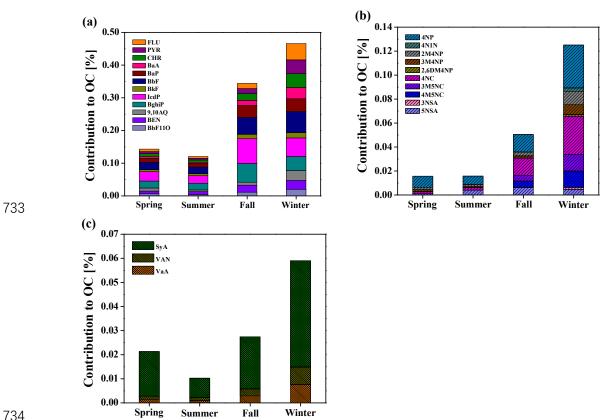
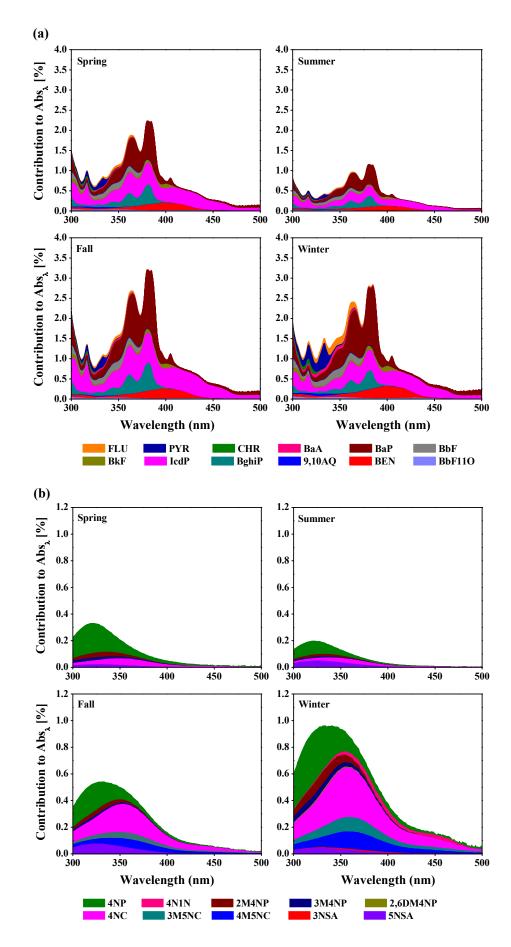


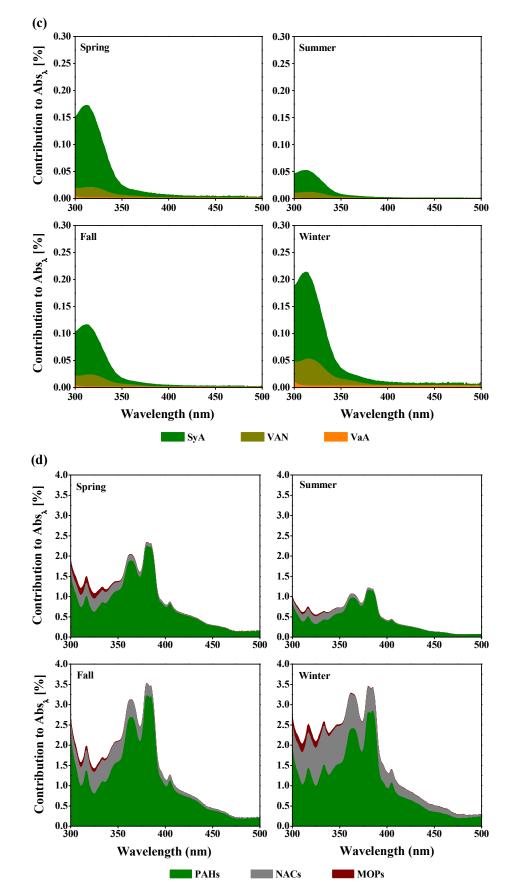
Figure 2. Comparison of AAE (left column) and MAE<sub>365</sub> (right column) values of water-soluble
BrC at remote sites (Srinivas and Sarin, 2013; Bosch et al., 2014; Zhang et al., 2017b), rural
sites (Hocobian et al., 2010; Kirillova et al., 2014a; Zhu et al., 2018; Xie et al., 2019) and urban
sites (Kirillova et al., 2014b; Yan et al., 2015; Chen et al., 2018; Huang et al., 2018; Park et al.,
2018).

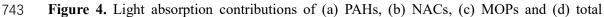




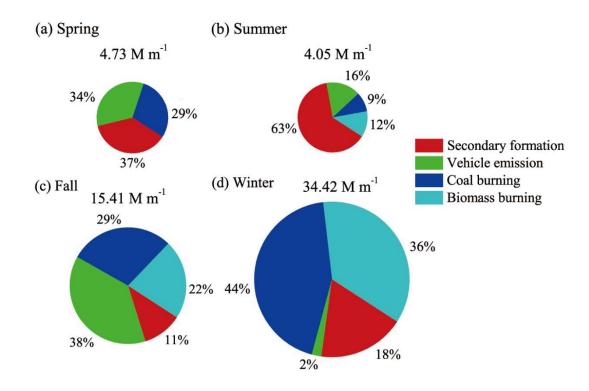
- Figure 3. Contributions of (a) PAHs, (b) NACs, and (c) MOPs carbon mass concentrations to
- the total OC concentrations.







- 744 measured chromophores to Abs<sub>MSOC</sub> over the wavelength range of 300 to 500 nm in spring,
- summer, fall and winter.





**Figure 5.** Contributions of the major sources to Abs<sub>365,MSOC</sub> in Xi'an during spring, summer, fall

and winter.