

# Evaluating Trends and Seasonality in Modeled PM<sub>2.5</sub> Concentrations Using Empirical Mode Decomposition

Huiying Luo<sup>1</sup>, Marina Astitha<sup>1\*</sup>, Christian Hogrefe<sup>2</sup>, Rohit Mathur<sup>2</sup>, S. Trivikrama Rao<sup>1,3</sup>

<sup>1</sup>University of Connecticut, Department of Civil and Environmental Engineering, Storrs-Mansfield, CT, USA

<sup>2</sup>U.S. Environmental Protection Agency, Research Triangle Park, NC, USA

<sup>3</sup>North Carolina State University, Raleigh, NC, USA

\*Corresponding author: Marina Astitha, Civil and Environmental Engineering, University of Connecticut, 261 Glenbrook Road, Storrs, CT, 06269-3037, Phone: 860-486-3941, Fax: 860-486-2298, Email: [marina.astitha@uconn.edu](mailto:marina.astitha@uconn.edu).

**Abstract.** Regional-scale air quality models are being used for studying the sources, composition, transport, transformation, and deposition of fine particulate matter (PM<sub>2.5</sub>). The availability of decadal air quality simulations provides a unique opportunity to explore sophisticated model evaluation techniques rather than relying solely on traditional operational evaluations. In this study, we propose a new approach for process-based model evaluation of speciated PM<sub>2.5</sub> using improved Complete Ensemble Empirical Mode Decomposition with Adaptive Noise (improved CEEMDAN) to assess how well version 5.0.2 of the coupled Weather Research and Forecasting model - Community Multiscale Air Quality model (WRF-CMAQ) simulates the time-dependent long-term trend and cyclical variations in the daily average PM<sub>2.5</sub> and its species, including sulfate (SO<sub>4</sub>), nitrate (NO<sub>3</sub>), ammonium (NH<sub>4</sub>), chloride (Cl), organic carbon (OC), and elemental carbon (EC). The utility of the proposed approach for model evaluation is demonstrated using PM<sub>2.5</sub> data at three monitoring locations. At these locations, the model is generally more capable of simulating the rate of change in the long-term trend component than its absolute magnitude. Amplitudes of the sub-seasonal and annual cycles of total PM<sub>2.5</sub>, SO<sub>4</sub>, and OC are well reproduced. However, the time-dependent phase difference in the annual cycles for total PM<sub>2.5</sub>, OC, and EC reveal a phase shift of up to half year, indicating the need for proper temporal allocation of emissions and for updating the treatment of organic aerosols compared to the model version used for this set of simulations. Evaluation of sub-seasonal and inter-annual variations indicates that CMAQ is more capable of replicating the sub-seasonal cycles than inter-annual variations in magnitude and phase.

## Keywords

Model evaluation, coupled WRF-CMAQ, improved Complete Ensemble Empirical Mode Decomposition (EMD) with Adaptive Noise, Speciated PM<sub>2.5</sub>, Scale Separation, Seasonality, Trend

## 30 **1 Introduction**

31 It is well recognized that inhalable fine particulate matter (PM<sub>2.5</sub>) adversely impacts human health and the  
32 environment. Regional-scale air quality models are being used in health impact studies and decision-making related  
33 to PM<sub>2.5</sub>. Long-term model simulations of PM<sub>2.5</sub> concentrations using regional air quality models are essential to  
34 identify long-term trends and cyclical variations such as annual cycles in areas larger than what is covered by in-situ  
35 measurements. However, total PM<sub>2.5</sub> concentrations are challenging to predict because of the dependence on the  
36 contributions from individual PM<sub>2.5</sub> components, such as sulfates, nitrates, carbonaceous species, and crustal elements.  
37 In this context, a detailed process-based evaluation of the simulated speciated PM<sub>2.5</sub> must be carried out to ensure  
38 acceptable replication of observations so model users can have confidence in using regional air quality models for  
39 policy-making. Furthermore, process-based information can be useful for making improvements to the model.

40 Some of the trend or step change evaluations of regional air quality models in the past have focused on specific pairs  
41 of years (Kang et al., 2013; Zhou et al., 2013; Foley et al., 2015). These studies do not properly account for the sub-  
42 seasonal and inter-annual variations between those specific periods. Trend evaluation is commonly done by linear  
43 regression of indexes such as the annual mean or specific percentiles, assuming linearity and stationarity of time series  
44 (Civerolo et al., 2010; Hogrefe et al., 2011; Banzhaf et al., 2015; Astitha et al., 2017). The problem with the linear  
45 trend evaluation is that there is no guarantee the trend is actually linear during the period of the study because the  
46 underlying processes are in fact nonlinear and nonstationary (Wu et al., 2007).

47 Seasonal variations are usually studied and evaluated by investigating the monthly or seasonal means of total and/or  
48 speciated PM<sub>2.5</sub> (Civerolo et al., 2010; Banzhaf et al., 2015; Yahya et al., 2016; Henneman et al., 2017). Evaluation of  
49 ten-year averaged monthly mean (i.e., ten-year averaged mean in Jan., ..., Dec.) of PM<sub>2.5</sub> simulated with WRF/Chem  
50 against the Interagency Monitoring of Protected Visual Environments (IMPROVE) by Yahya et al. (2016) shows that  
51 the model captures the observed features of summer peaks in PM<sub>2.5</sub> with a phase shift of few months. However,  
52 according to the analysis (Fig. 10) in Henneman et al. (2017), the seasonality shown in monthly-averaged PM<sub>2.5</sub> time  
53 series is much less distinguishable compared with that of ozone and CMAQ (version 5.0.2) does not replicate the  
54 monthly PM<sub>2.5</sub> quite well with large underestimation in the summer months. In these studies, the seasonality might not  
55 be well represented by the preselected averaging window size of one or three months. In addition, averaging of those  
56 monthly or seasonal means across multiple years may conceal the long-term trends or interannual variations driven  
57 by climate change, emission control policies or other slow varying processes.

58 To address the above-mentioned problems, we propose a new method for conducting air quality model evaluation for  
59 PM<sub>2.5</sub> using improved CEEMDAN. Improved CEEMDAN is an Empirical Mode Decomposition (EMD)-based, data-  
60 driven intrinsic mode decomposition technique that can adaptively and recursively decompose a nonlinear and  
61 nonstationary signal into multiple modes called intrinsic mode functions (IMFs) and a residual (trend component)  
62 (Huang et al., 1998; Wu and Huang, 2009; Yeh et al., 2010; Torres et al., 2011; Colominas et al., 2014). It does not  
63 require any preselection of the temporal scales or assumptions of linearity and stationarity for the data, thereby  
64 providing some insights into time series of PM<sub>2.5</sub> concentrations and its components. Decomposed PM<sub>2.5</sub> long-term  
65 trend components and annual cycles from observed and simulated PM<sub>2.5</sub> serve as the intuitive carrier of the trend and

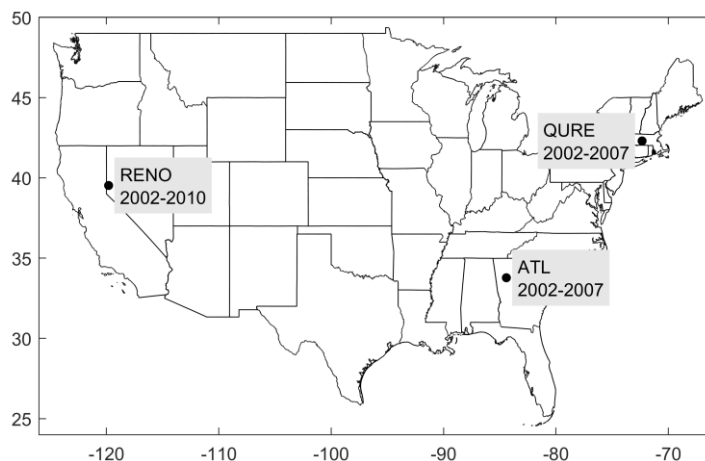
66 seasonality evaluation. In the meantime, several other IMFs with characteristic time scales ranging from multiple days  
67 to years are also decomposed, enabling model evaluation of the less studied sub-seasonal and inter-annual variations.

## 68 **2 Coupled WRF-CMAQ PM<sub>2.5</sub> Simulations and Observations**

69 The two-way coupled WRF-CMAQ (version 5.0.2) is configured with a 36 km horizontal grid spacing over the  
70 contiguous United States (CONUS) with 35 vertical layers of varying thickness extending from the surface to 50 mb  
71 (Wong et al., 2012; Gan et al., 2015). Time-varying chemical lateral boundary conditions were derived from the 108  
72 km resolution hemispheric WRF-CMAQ (Mathur et al., 2017) simulation for the 1990-2010 period (Xing et al., 2015).  
73 The simulations are driven by a comprehensive emission dataset which includes aerosol precursors and primary  
74 particulate matter (Xing et al., 2013, 2015). Annual emissions for the CMAQ simulations were estimated using the  
75 methodology described in Xing et al. (2013). Briefly, the National Emissions Inventory (NEI) for 1990, 1995, 1996,  
76 1999, 2001, 2002, and 2005 and a number of sector-specific long-term databases containing information about trends  
77 in activity data and emission controls were used to create county-level annual emissions for a total of 49 emission  
78 sectors. Prior to being used as input to the CMAQ simulations, these annual emissions were then temporally and  
79 spatially allocated to provide hourly emissions based on monthly, weekly, and diurnal temporal cross-reference and  
80 profile data from the 2005 NEI modeling platform. These profile data vary by emission sources and sometimes by  
81 state and county and are generally based on surveys and extrapolation of activity data which can be subject to  
82 uncertainty. Exceptions to the use of 2005 NEI platform temporal profile data for temporal allocation were emissions  
83 from electric generating units (EGU) which directly used measured hourly emissions after 1995 and wildfire emissions  
84 that used climatological monthly, weekly, and diurnal profiles for temporal allocation. Readers can refer to Gan et al.  
85 (2015) for additional model information and the trend evaluation against seven pairs of sites from the CASTNET  
86 (Clean Air Status and Trend Network) and IMPROVE networks for 1995-2010. We obtained the 2002-2010 daily  
87 average PM<sub>2.5</sub> and its speciated time series from the set of simulations with direct aerosol feedback. The earlier years  
88 of 1990-2001 are not included in this evaluation because of the limited availability of speciated PM<sub>2.5</sub> observations.

89 To avoid misinterpretation of data due to the presence of missing values, only sites with continuous complete long-  
90 term record for total PM<sub>2.5</sub> and its speciation including SO<sub>4</sub>, NO<sub>3</sub>, NH<sub>4</sub>, OC, EC, and Cl are studied (Fig. 1). All of the  
91 selected sites have data coverage above 90% each year for at least six consecutive years between 2002 and 2010  
92 (equivalent to 30% for 1-in-3 days sampling sites). This strict data selection led to the sparsity of this type of  
93 observations for the study period. QURE, a rural site carrying out 1-in-3 days sampling of total and speciated PM<sub>2.5</sub>  
94 of SO<sub>4</sub>, NO<sub>3</sub>, OC, EC, and Cl, is located in Quabbin Summit, MA. It is one of the three sites from the IMPROVE  
95 network that has at least six continuous years of speciated observations and was selected here to demonstrate the  
96 application of the proposed method in rural areas. It should be noted that the majority of the observed Cl in 2002 and  
97 2003 is negative due to a filter issue which was not addressed until 2004 (White, 2008). Thus, simulations of Cl are  
98 only evaluated during 2004-2007 at this site. Station RENO, located in urban Reno, NV, is also a 1-in-3 days sampling  
99 site of total and speciated PM<sub>2.5</sub> of SO<sub>4</sub>, NO<sub>3</sub>, NH<sub>4</sub>, OC, and EC, and it is the only Chemical Speciation Network  
100 (CSN) site that fulfills this data coverage requirement. The third site ATL in the Southeastern Aerosol Research and  
101 Characterization Study (SEARCH) network is located 4.2 km northwest of downtown Atlanta, GA. It is the only long-

102 term site available with daily sampling (Hansen et al., 2003; Edgerton et al., 2005) that meets the data coverage  
103 requirement. The best-estimate (BE), a calculated concentration intended to represent what is actually in the  
104 atmosphere (Edgerton et al., 2005), of the total  $PM_{2.5}$  and  $SO_4$ ,  $NO_3$ ,  $NH_4$ , and EC components are retrieved for the  
105 evaluation. OC component is a direct measurement. These three sites have a continuous record covering at least 6  
106 years (2002 – 2007 for QURE and ATL, 2002 – 2010 for RENO) that allows an evaluation of long-term trends.



107  
108 **Fig. 1. Location and data coverage of the  $PM_{2.5}$  monitoring sites QURE, RENO, and ATL.**

### 109 **3 Methodology**

#### 110 **3.1 Empirical Mode Decomposition**

111 The Empirical Mode Decomposition (EMD) technique, proposed in the late 1990s, is capable of adaptively and  
112 recursively decomposing a signal into multiple modes called intrinsic mode functions (IMFs), where each mode has  
113 a characteristic frequency, and a residual with one extremum at most (Huang et al., 1998). EMD decomposes the  
114 original signal into several IMFs and a residual through a repeated process called “sifting”: first, local maxima and  
115 minima are identified and interpolated separately with a cubic spline as the upper and lower envelop; then an IMF  
116 candidate is derived by subtracting the mean of the envelopes from the original signal. If the candidate satisfies the  
117 following criteria (Huang et al., 1998), it is saved as the first IMF (IMF1), and the remaining portion (original signal  
118 – IMF1) is treated as a new input signal for the decomposition of the remaining IMFs; otherwise, more sifting  
119 processes should be carried out until the candidate becomes an IMF.

120 1) The number of extrema (maxima and minima) and the number of zero-crossings must be equal or differ at most by  
121 one; 2) The local mean at any point, the mean of the envelope defined by local maxima and the envelope defined by  
122 local minima, must be zero.

123 In this way, IMF1, IMF2, ... are decomposed recursively with decreasing characteristic frequency. The final remaining  
124 residual (trend) could be a monotonic function of time or a long-term component with one extremum at most. The  
125 decomposed signal then is expressed as the summation of all IMFs and the final residual:

126 
$$x = \sum_{i=1}^k d_i + r \quad (1)$$

127 where  $x$  is the original signal,  $d_i$  is the  $i^{\text{th}}$  IMF,  $k$  is the total number of IMFs and  $r$  is the final residual.

128 Nevertheless, “mode mixing”, where oscillations with very disparate scales can be present in one mode or vice versa,  
129 is commonly reported. To cope with this issue, multiple noise assisted EMD have been developed successively (Wu  
130 and Huang, 2009; Yeh et al., 2010; Torres et al., 2011; Colominas et al., 2014). It is evident that the latest improved  
131 Complete Ensemble EMD with Adaptive Noise (improved CEEMDAN) manages to alleviate the problem of mode  
132 mixing with the benefit of reducing the amount of noise presented and avoiding spurious modes (Colominas et al.,  
133 2014). Moreover, the end effects or boundary effects have been addressed by its predecessor EEMD (Ensemble  
134 Empirical Mode Decomposition) by extrapolating the maxima and minima, and behaved well in numerous time series  
135 with dramatically variant characteristics (Wu and Huang, 2009). The extrapolation of maxima and minima is proven  
136 to be more effective compared with the extrapolation of the signal itself such as repetition or reflection (Rato et al.,  
137 2008).

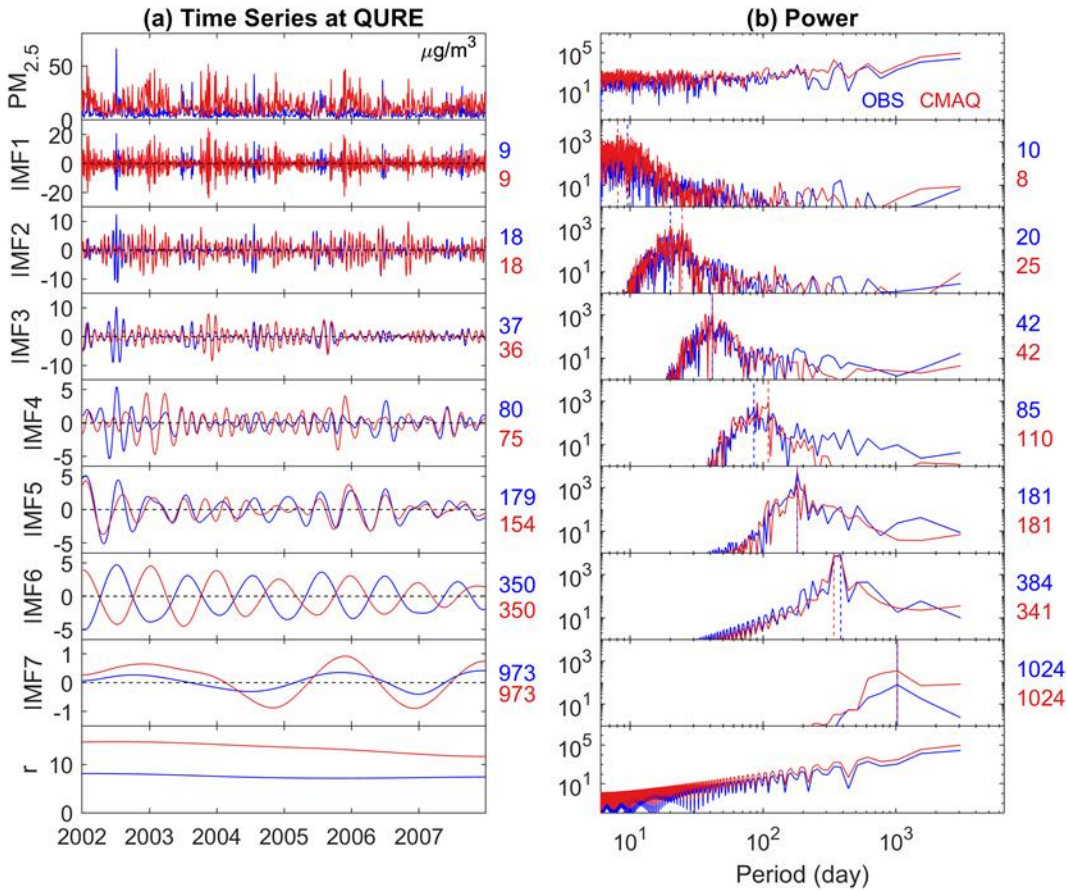
138 Given the EMD’s ability to deal with real-world nonstationary and nonlinear time series, it is widely used in  
139 engineering, economics, earth, and environmental sciences (e.g., Huang et al., 1998; Chang et al., 2003; Yu et al.,  
140 2008; Colominas et al., 2014; Derot et al., 2016). We use the most up-to-date noise-assisted improved CEEMDAN  
141 technique with at least hundreds of noise realizations to decompose observed and simulated PM<sub>2.5</sub> time series. Readers  
142 can refer to Colominas et al. (2014) for a detailed description of the technique and access to the corresponding  
143 MATLAB code. Trial and error attempts are made in setting the inputs (standard deviation of the added noise and the  
144 limit of maximum number of sifting allowed) of the improved CEEMDAN function to achieve best mode separation.  
145 In a desired best mode separation, neighboring IMFs should have very limited levels of mode mixing, which can be  
146 fast screened based on the time series of the decomposed IMFs and their power spectrum.

147 The impact of boundaries on the decomposed annual cycles and the residual is assessed by the variations (standard  
148 deviation) of hypothetical decomposed boundaries by cutting a continuous eighteen-year total PM<sub>2.5</sub> observation  
149 (North Little Rock, AR) 48 times at different years and times of the year (Fig. S1). The standard deviation of the  
150 annual cycles is found to largely diminish within half the annual cycles and could be negligible within one year. This  
151 could very possibly expand to IMFs with other characteristic scales. Yet, trend components (residuals) show  
152 variability depending on the available time period after cutting. Most of the time, they follow the reference long-term  
153 trend reflected either by the residual or the summation of the residual and the IMF with the longest temporal scale  
154 decomposed from the eighteen-year PM<sub>2.5</sub> (Fig. S1c). This is in line with our expectations as a trend should exist  
155 within a given time span, following the definition in Wu et al. (2007): “The trend is an intrinsically fitted monotonic  
156 function or a function in which there can be at most one extremum within a given data span”.

157 Although very strict data completeness requirement is employed for this study, it should not be conceived as a  
158 limitation of the method itself. A sensitivity test based on a period of nine years of total PM<sub>2.5</sub> observations at the same  
159 site with 99% data coverage shows that even though variability of annual cycles and long-term trends increases with  
160 decreased data availability (100%, 90%..., 10%), the structure of those components is consistent. The average of 40  
161 realizations of annual cycles and long-term trend components in each data-completeness scenario is in perfect

162 alignment with that of 100% data completeness (Fig. S2 and S3). Given the fact that those 40 realizations in each  
 163 scenario are based on independent random samplings of the original observations, the increased variability could very  
 164 possibly result from the difference in the sampled data itself rather than the method. Thus, the robustness of improved  
 165 CEEMDAN decomposed annual cycles and long-term trend is justified. In fact, EMD has been proven to be an  
 166 effective tool for data gap-filling (Moghtaderi et al., 2012).

167



168

169 **Fig. 2.** Decomposition of observed (blue) and simulated (red) 24-hour average total PM<sub>2.5</sub> into 7 IMFs and a  
 170 residual component (trend) at Quabbin Summit, MA using the improved CEEMDAN: (a) Time series of total  
 171 PM<sub>2.5</sub>, IMFs, and the residual component (all with the unit of µg/m<sup>3</sup>); (b) Power spectrum of the corresponding  
 172 time series. The colored numbers on the right side of time series are the mean period  $t_m$  in days, while the ones  
 173 on the right side of the power spectrum are the peak period  $t_p$  in days, which are also indicated by the dashed  
 174 vertical lines on the power spectrum. Note that the scales for the time series are not all the same. Also, all power  
 175 spectra are in the log scale, and those of the IMFs are zoomed in with a range of 10<sup>0</sup> to 10<sup>4</sup> on the y-scale for  
 176 better visual clarity (compared with 10<sup>-2</sup> to 10<sup>7</sup> for total PM<sub>2.5</sub> and the residual component).

177

178 The characteristic period of each IMF can be estimated by the peak period  $t_p$  (days) where the power spectrum of the  
 179 IMF peaks:



180 
$$t_p = \frac{1}{f_p} \quad (2)$$

181 in which  $f_p$  is the frequency that the power spectrum peaks in the unit of number of cycles per day. The peak estimates  
 182 can be biased if more than one high-power frequency is located closely within one IMF. Thus, the power spectrum  
 183 and  $t_p$  is only used as a fast screening tool to determine if a desired decomposition is accomplished. As an alternative  
 184 approach, the mean period  $t_m$  can be estimated by:

185 
$$t_m = \frac{\text{Time span}}{(n_{max} + n_{min} + n_{zero})/4} \quad (3)$$

186 where  $n_{max}$ ,  $n_{min}$ , and  $n_{zero}$  are the number of maxima, minima, and zero-crossings, respectively, during the  
 187 *Time span* (days). As the frequency decreases, the mean period estimates become less accurate because of the limited  
 188 time span compared with the length of the cycle and should be carefully interpreted.

189 An example of the total PM<sub>2.5</sub> decomposition with improved CEEMDAN at the QURE site shows modes ranging from  
 190 very high frequency to very low frequency (IMF1 to IMF7) and a residual (Fig. 2). No visible mode mixing can be  
 191 detected in both the time series (Fig. 2a) and the power spectrum (Fig. 2b) of all IMFs. Mean ( $t_m$ ) and peak ( $t_p$ )  
 192 estimations of the characteristic periods of each IMF are presented on the right side of each mode. Annual cycles and  
 193 long-term trend components are well represented by IMF6 and the residual, with the remaining IMFs carrying weekly,  
 194 sub-seasonal, seasonal, and inter-annual variations, respectively, for both observed and simulated PM<sub>2.5</sub> (Fig. 2). We  
 195 have noticed that in some rare cases, a spurious mode in the last IMF with synchronous signal and very close scales  
 196 to its previous IMF exists. This is possibly due to the fact that the characteristic periods of those IMFs are in proximity  
 197 to the span of the studied time span. In these cases, the last two modes are merged by adding them together to conduct  
 198 a detailed evaluation as discussed in Section 4.1.

### 199 3.2 Statistical metrics

200 EMD-decomposed IMFs and trend components allow for a detailed time-dependent evaluation of PM<sub>2.5</sub> and provide  
 201 a novel opportunity to trace the performances of specific scales back to the corresponding speciated components. Note  
 202 that the trend component is the decomposed residual component from the PM<sub>2.5</sub> in the unit of  $\mu\text{g}/\text{m}^3$ , and it is not the  
 203 traditional concept of trend in concentration per time. In addition to a direct evaluation of its magnitude, we also  
 204 calculated its derivative to identify the periods with higher or lower rate of change (concentration per time). Time-  
 205 dependent intrinsic correlation (TDIC) is utilized to study the evolvement of the model performance for cyclic  
 206 variations throughout time (Chen et al., 2010; Huang and Schmitt, 2014; Derot et al., 2016). It is a set of correlations  
 207 calculated for IMFs over a local period of time  $I$  centered around time  $t$ :

208 
$$I(t) = [t - \frac{t_w}{2}, t + \frac{t_w}{2}] \quad (4)$$

209 in which  $t$  is the center time for the calculation of the correlation and  $t_w$  is the moving window length. The minimum  
 210 of  $t_w$  is set to be the local instantaneous period of the IMF (larger of that in observation or simulation) using the  
 211 general zero-crossing method to ensure that at least one instantaneous period is included in calculating the local

212 correlation coefficient (Chen et al., 2010). The maximum of  $t_w$  is the entire data period with a traditional overall  
 213 correlation being calculated. The empty spaces in the pyramids used to depict the TDIC are an indication that the  
 214 correlation is not statistically significantly different from zero. With both decomposed observed and modeled  
 215 concentrations in a narrow scale range, the correlation would no longer be contaminated by coexisting signals of  
 216 different scales (Chen et al., 2010).

217 In order to summarize the performance of the decomposed trend components and IMFs, the ratio of the mean  
 218 magnitudes of the trend components is defined as:

$$219 \quad r_{trend} = \frac{Mean_{CMAQ}}{Mean_{observation}} \quad (5)$$

220 where  $Mean_{CMAQ}$  and  $Mean_{observation}$  represent the mean of simulated and observed residual components  
 221 respectively. The ratio of the mean amplitude of each IMF is defined by Equation 6, where an example for the annual  
 222 cycles is provided:

$$223 \quad r_{annual} = \frac{RMS_{CMAQ,annual}}{RMS_{observation,annual}} \quad (6)$$

224 where  $RMS_{observation,annual}$  and  $RMS_{CMAQ,annual}$  represent the root mean square of observed and simulated annual  
 225 cycles respectively. Finally, the phase shift of an IMF  $n$  is defined as the days an IMF decomposed from modeled  
 226 time series has to be shifted to maximize the correlation ( $R_{max}$ ) with the corresponding IMF from observed PM<sub>2.5</sub>  
 227 time series. In practice,  $n$  could be as much as a few cycles of the mean period,  $t_m$ . Here, we limit the absolute number  
 228 of shift days to not exceed a half cycle as a reference for the phase shift of an IMF. Thus,  $n$  satisfies  $-(t_m/2) \leq n \leq$   
 229  $(t_m/2)$  with  $t_m$  being the larger mean period in observation or simulation. It becomes  $-0.5 \leq n/t_m \leq 0.5$  in terms  
 230 of number of cycles.

## 231 **4 Results and Discussion**

### 232 **4.1 Temporal scales**

233 Temporal scales in PM<sub>2.5</sub> resolved by EMD depend solely on the intrinsic properties of the data itself. These properties  
 234 include underlying characteristics of specific PM<sub>2.5</sub> concentrations, the data sampling frequency, which determines the  
 235 scales that can be resolved in the high frequency IMFs, and the time span for the data coverage, which could possibly  
 236 play an important role in differentiating the low frequency IMFs from the trend component. Here, we first evaluate  
 237 the scales represented by the mean period in the speciated and total PM<sub>2.5</sub> time series. Since each IMF represents a  
 238 nonstationary process, the mean period  $t_m$  is only an estimate of its characteristic scale. Evaluation of  $t_m$  might not  
 239 necessarily be able to identify issues with corresponding model simulations, and it does not indicate any information  
 240 on the magnitude or the phase of the time series, which is more important and will be further discussed in Sections  
 241 4.3 to 4.4.

242 Fig. 3a presents the characteristic scales ( $t_m$ ) of IMFs in observed and simulated total and speciated PM<sub>2.5</sub> of QURE.  
 243 The CMAQ model compares well with the observations for IMFs 1 through 6 with cycles of 9, 19, 37, 78, 158 and



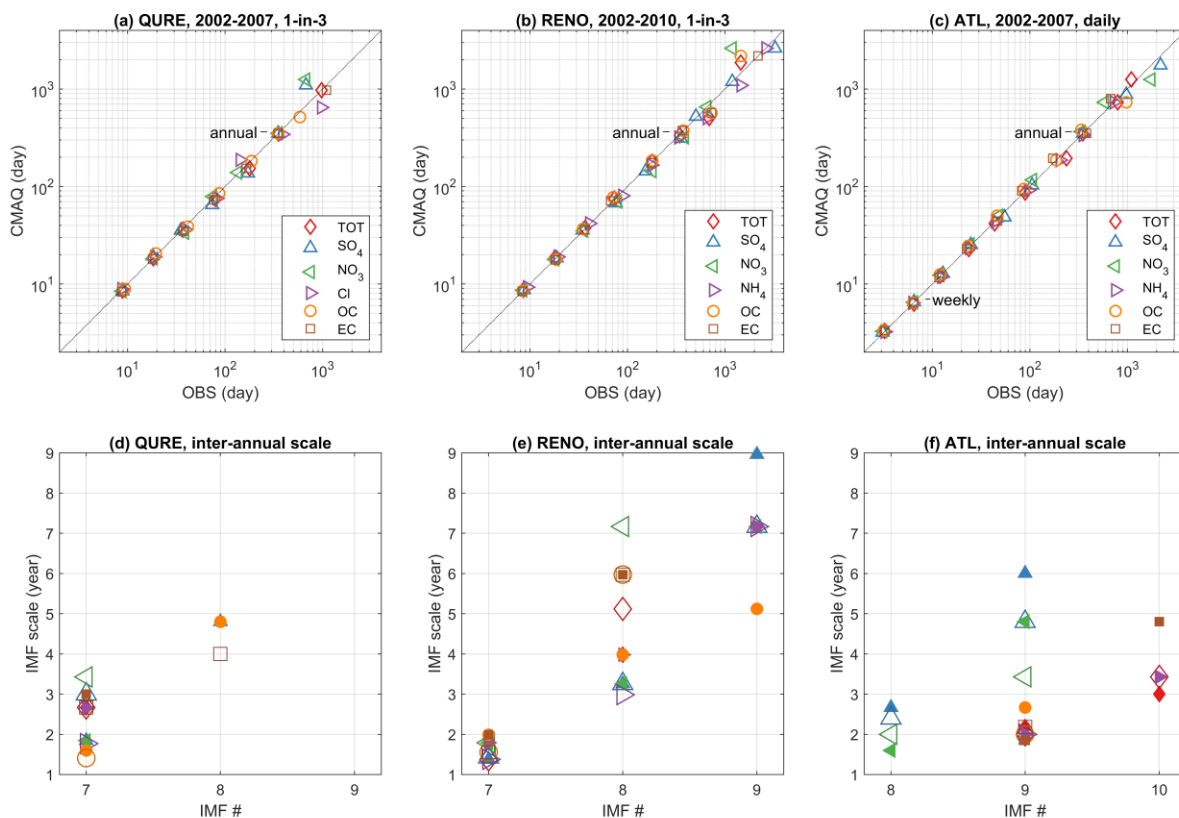
244 347 days (average of all observed and simulated total and speciated  $PM_{2.5}$ ). Among all these IMFs, IMF6, which  
245 represents the annual cycles, shows the least variations in the characteristic scale (Fig. 3a) and highest peak energy  
246 from the power spectrum such as Fig. 2b for total  $PM_{2.5}$ , except for observed EC and OC where the power of half-  
247 year cycles is more dominant (Fig. S4). These two features demonstrate a clear seasonality in both observed and  
248 simulated total and speciated  $PM_{2.5}$ , which would otherwise be concealed by practices such as monthly averaging.  
249 This can be further confirmed by the statistical significance of the annual cycles (except for observed EC and OC)  
250 (Fig. S5) based on a Monte Carlo verified relationship between the energy density and mean period of IMFs (Wu and  
251 Huang, 2004; Wu et al., 2007). To explore the inter-annual cycles in more detail, mean periods of IMFs with scales  
252 longer than a year are being displayed in the top left panel of Fig. 3a. Some variability exists between the observation  
253 and model simulation to the extent that not all IMFs from observation are being simulated and vice versa for the inter-  
254 annual cycles. The characteristic scales of all decomposed IMFs with scales longer than a year are shown in Fig. 3d.  
255 The estimated mean periods of the inter-annual cycles and the differences in the presence of slow varying cycles with  
256 the long characteristic scales are likely to be influenced by their proximity to the data time span of 6 years (4 years for  
257 Cl). This implies that the model evaluation shouldn't go beyond 3 years (2 years for Cl) given the current data  
258 coverage. CMAQ captured the 3-year cycles in EC and total  $PM_{2.5}$  and 2-year cycles in OC and Cl, despite an  
259 overestimation in the scales of 2-year cycles in observed  $SO_4$  and  $NO_3$ .

260

261

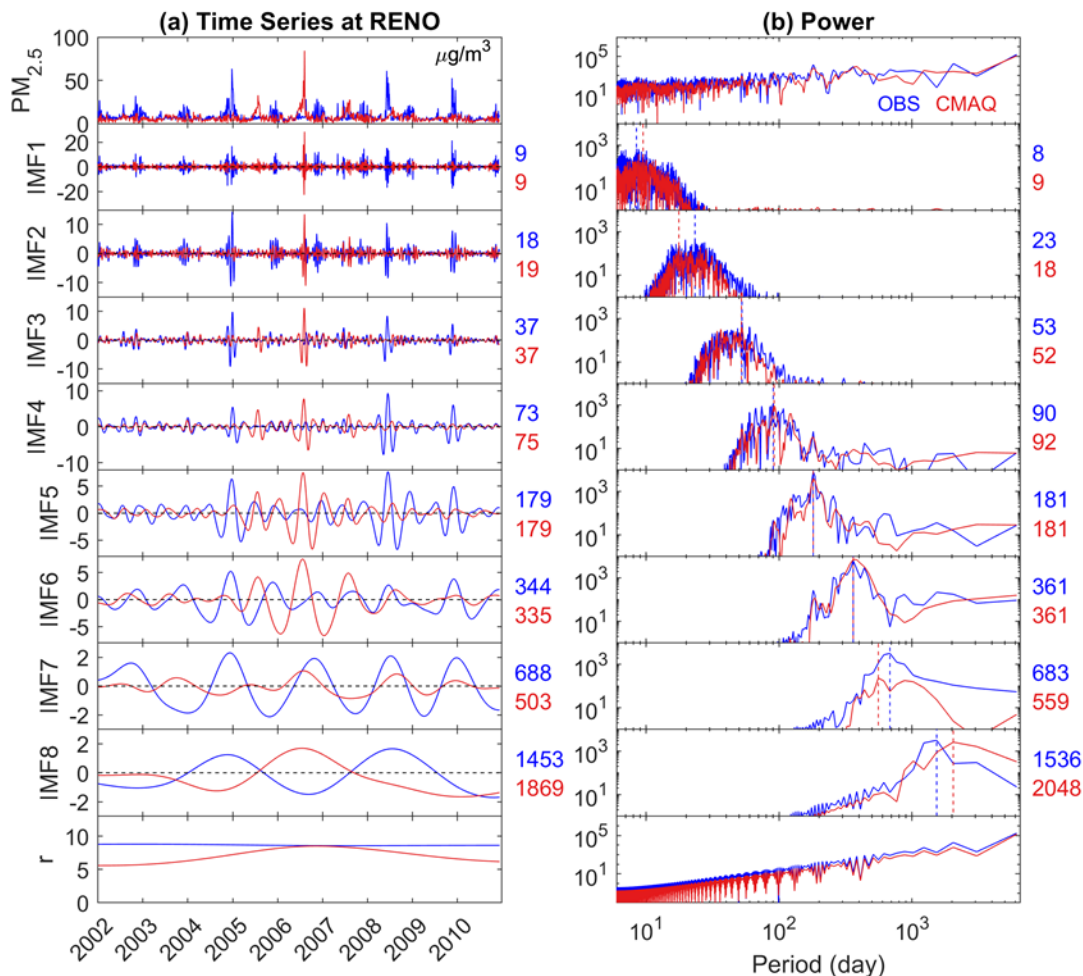
262

263



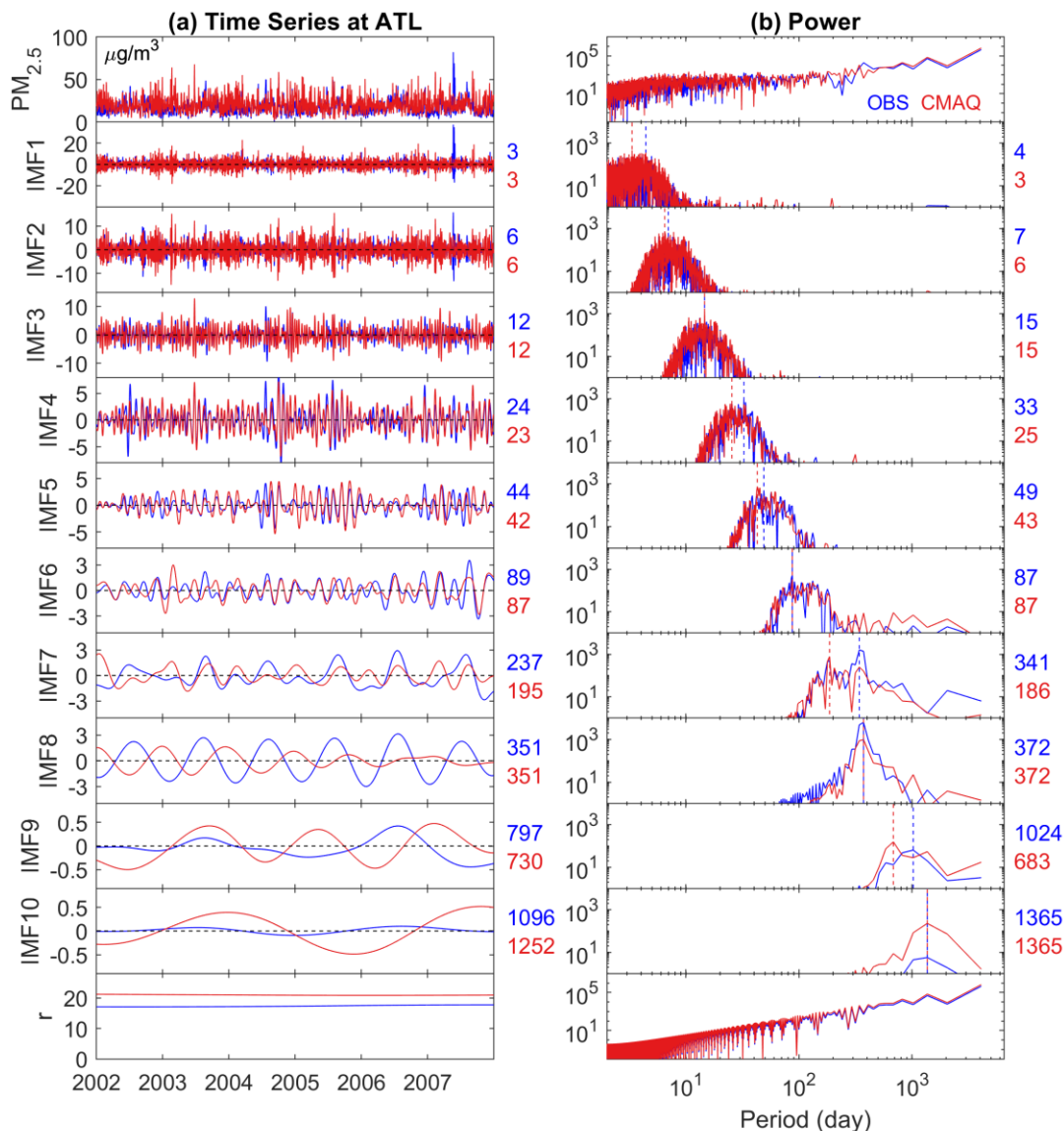
264  
 265 **Fig. 3. The characteristic scales ( $t_m$ ) resolved in the IMFs of observed and simulated total and speciated  $PM_{2.5}$**   
 266 **for (a, d) QURE, (b, e) RENO, and (c, f) ATL. In (a-c), IMF1 to the last pair of IMFs with increasing**  
 267 **characteristic periods are shown from bottom left to top right. Mean periods of IMFs with scales longer than a**  
 268 **year are being displayed in (d-f) with the same shapes as in the legend above to show the characteristic scales**  
 269 **of all decomposed IMFs given that not all IMFs from observation are being simulated and vice versa. In the (d-**  
 270 **f), species decomposed from observations are shown with smaller filled shapes, while species decomposed from**  
 271 **simulations are represented by larger open shapes in slightly darker shades.**

272 Similar features in observed and simulated total and speciated  $PM_{2.5}$  concentrations at RENO are presented in Fig. 3b.  
 273 Likewise, the highest peaks in the power spectrum also sit in the annual cycles of IMF6 except for the observed OC  
 274 and total  $PM_{2.5}$  which have higher peak power at half-year cycles. All annual IMFs are statistically significant except  
 275 for simulated  $NH_4$  (Fig. S5). The small variation in the estimated characteristic period of IMF6 is because this  
 276 monitoring site is located in a wildfire prone region on the border of Nevada and California. Clear evidence can be  
 277 seen from Fig. 4a that an extra annual cycle in the IMF6 of observations in the summer of 2008 is depicted, which is  
 278 very possibly driven by the 2008 California Wildfires spanning from May until November. Satellite image of the  
 279 wildfire smoke on July 10, 2008 can be found in Figure 1 from Gyawali et al. (2009). Unlike the diversified scales in  
 280 IMF7 at QURE, IMF7 at RENO features universal 2-year cycles of all species as well as total  $PM_{2.5}$  and all of them  
 281 are well replicated by the model. However, variations in time scales are present in IMF8 possibly because of the  
 282 limited data coverage. Thus, only species with time scales less than 4 years in both observations and model simulations  
 283 are evaluated. It is evident that CMAQ has reproduced the 3-year cycles in  $SO_4$  and  $NH_4$ .



284  
 285 **Fig. 4. Same as Fig. 2 but for the RENO site with 8 IMFs.**  
 286

287 ATL is the only speciated site with daily data coverage. Observed and simulated total and speciated PM<sub>2.5</sub>  
 288 concentrations at the ATL site are decomposed into 9 or 10 IMFs (Fig. 3c). Because of the change in data frequency,  
 289 high frequency scales such as weekly cycles can be evaluated and the significance tested (Fig. S5) annual cycles with  
 290 the highest peak power is represented by IMF8 (IMF7 for SO<sub>4</sub> and NO<sub>3</sub>). Annual cycles of SO<sub>4</sub> and NO<sub>3</sub> appeared in  
 291 the earlier stage of decomposition in IMF7 because of their relatively weak half-year cycles, which largely led to the  
 292 mixed signal of half-year and annual cycles in IMF7 in total PM<sub>2.5</sub> as in Fig. 5b. This is more visible in the observed  
 293 IMF7 where the energy of the one-year period surpasses that of the half-year. Yet, clues can be seen from Fig. 5 that  
 294 the amplitude and the energy of annual cycles leaked into IMF7 is very limited compared to that remaining in IMF8,  
 295 indicating that it is still safe to conduct model evaluation on the seasonality using IMF8 with an underestimation in  
 296 the amplitude of observation. On the other hand, inferences should be made with caution for IMF7 because of the  
 297 mixed modes. Scales up to 3 years are relatively well reproduced by the model.



298

299 **Fig. 5. Same as Fig. 2 but for the ATL site with 10 IMFs.**

### 300 **4.2 Long-term trend**

301 The EMD-decomposed long-term trend components for the observed and simulated total and speciated PM<sub>2.5</sub>

302 concentrations are presented in Fig. 6. To better visualize the non-linearity of the trend component,

303 the rates of change (temporal derivative of a trend component, which is the change in the consecutive concentration divided by the

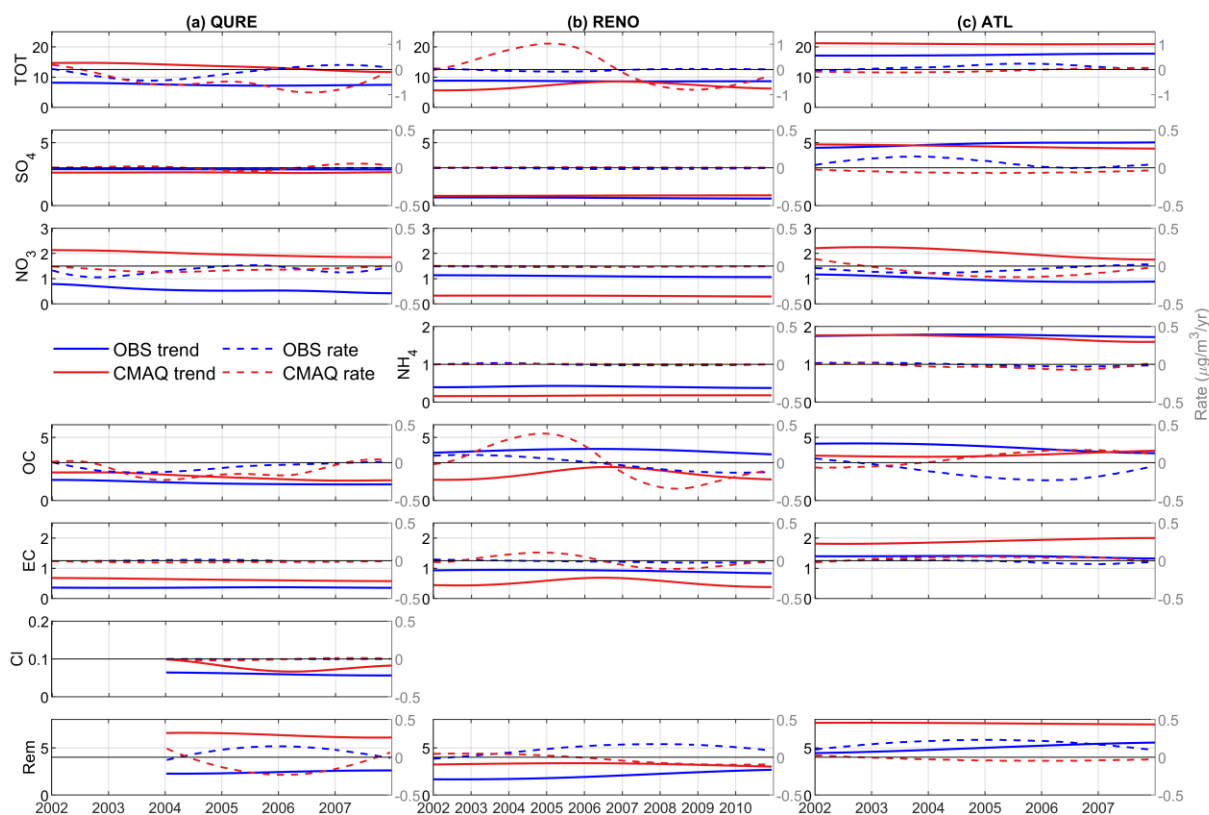
304 sampling rate of 1 or 3 days and converted to the unit of  $\mu\text{g}/\text{m}^3/\text{year}$  by multiplying 365 day/year) are added with a

305 separate y-axis on the right side in each panel (gray colored scale). It is evident that PM<sub>2.5</sub> is changing at a varying

306 rate, forming either a monotonic trend component or a trend component with one extremum, which cannot be fully

307 represented by a single constant number using a traditional linear regression approach. Given that there are chemical

308 species (the remaining component, *Rem*) other than the ones studied in the total PM<sub>2.5</sub>, not all performance issues can  
 309 be fully explained by the five available species.



310  
 311 **Fig. 6. Trend components of observed and simulated total and speciated PM<sub>2.5</sub> for (a) QURE, (b) RENO and (c)**  
 312 **ATL in  $\mu\text{g}/\text{m}^3$ . Dashed lines representing the rate of the change (temporal derivative of the trend component**  
 313 **converted to  $\mu\text{g}/\text{m}^3/\text{year}$ ) are plotted against the right-side y axis, with a reference line of no change in black in**  
 314 **the center. Note that the scales are not all the same.**

315  
 316 At the QURE site, CMAQ captures the general decreasing trend in observed total PM<sub>2.5</sub> which can mainly be traced  
 317 back to NO<sub>3</sub>, OC, and the remaining components, while both observed and simulated trend components of SO<sub>4</sub> and  
 318 EC are relatively constant (Fig. 6a). The relative importance of each component in driving the trend of observed and  
 319 simulated total PM<sub>2.5</sub> reflected by its mean concentration share is summarized in Table 1 (time-dependent variations  
 320 of the concentration share is attached in Fig. S6 for reference). Moreover, the periods with highest decreasing rate in  
 321 observed total PM<sub>2.5</sub> during 2003-2004 with a decreasing rate of  $-0.44 \mu\text{g}/\text{m}^3/\text{year}$  is also well replicated by the model.  
 322 Nevertheless, the slightly increasing PM<sub>2.5</sub> level in the later years is simulated to be decreasing at a much higher rate,  
 323 which is partly due to the overestimated decreasing rate in OC and species other than the five studied ones. The trend  
 324 component of simulated Cl shows a cyclic-like feature because of proximity between the existence of a cycle of 4-5

325 years (by decomposing the simulation during the 6-year study period) and 4-year period limited by the available  
 326 quality assured observations. The rate of change in the simulated trend component by decomposing the simulation  
 327 during the 6-year study period would mimic that from the 4-year observation, both with a negligible negative value  
 328 throughout 2004-2007. However, the mean magnitude of the trend component is almost twice as high (1.8 times  
 329 compared with observation) in the model with contribution from all species except for SO<sub>4</sub>. A quantitative summary  
 330 of the comparison between the mean magnitudes of the observed and model trend components can be found in Table  
 331 2.

332 **Table 1.** Concentration share (%) of different components in total PM<sub>2.5</sub>. It is estimated by dividing the mean trend  
 333 components of each species by that of total PM<sub>2.5</sub> for both OBS and CMAQ, multiplied by 100. The concentration  
 334 share of the remainder species *Rem* is estimated by subtracting all the available species share from 100 to compensate  
 335 for the small discrepancies caused by the rounding up process and uncertainty in the mode decomposition. “-” indicates  
 336 the data is not available (same applies for all other tables).

		SO <sub>4</sub>	NO <sub>3</sub>	NH <sub>4</sub>	OC	EC	Cl	Rem
<b>QURE</b>	OBS	38	7	-	19	5	1	30
	CMAQ	19	15	-	14	5	1	47
<b>RENO</b>	OBS	7	13	5	46	11	-	20
	CMAQ	11	4	2	30	7	-	45
<b>ATL</b>	OBS	28	6	10	24	8	-	24
	CMAQ	22	10	8	17	9	-	33

337

338 **Table 2.** The ratio of mean magnitude of the trend component  $r_{trend}$  (CMAQ/observation). Boldface values indicate  
 339 a relatively good estimate of the magnitude (0.7 - 1.3).

	TOT	SO <sub>4</sub>	NO <sub>3</sub>	NH <sub>4</sub>	OC	EC	Cl
QURE	1.8	<b>0.9</b>	3.5	-	1.4	1.7	<b>1.3</b>
RENO	<b>0.8</b>	<b>1.3</b>	0.3	0.4	0.5	0.6	-
ATL	<b>1.2</b>	<b>1.0</b>	2.1	<b>1.0</b>	<b>0.9</b>	1.4	-

340

341

342 RENO is located close to the border with California and is affected by large wildfire breakouts in the western U.S.  
 343 (Gyawali et al., 2009) as can be seen in the spikes of the observed total PM<sub>2.5</sub> (Fig. 4a). Thus, OC makes up a much  
 344 larger portion of total PM<sub>2.5</sub> compared to other locations (Table 1). The model simulates large increasing rate up to  
 345 1.03 µg/m<sup>3</sup>/year and decreasing rate up to -0.80 µg/m<sup>3</sup>/year before and after the 2006-2007 winter season and fails to  
 346 reproduce the relatively stable condition seen in the observations with only -0.09 µg/m<sup>3</sup>/year decreasing in 2004-2005  
 347 and 0.04 µg/m<sup>3</sup>/year increasing in 2008-2009 (Fig. 6b). Similar feature is found for combustion-related OC and EC  
 348 species. The observed slightly decreasing trends in SO<sub>4</sub> and NH<sub>4</sub> during 2005-2009 are not being captured in the model  
 349 simulations. The magnitude of the trend component is slightly underestimated with  $r_{trend}$  of 0.8 with contribution  
 350 from all species except for SO<sub>4</sub> as well (Table 2).

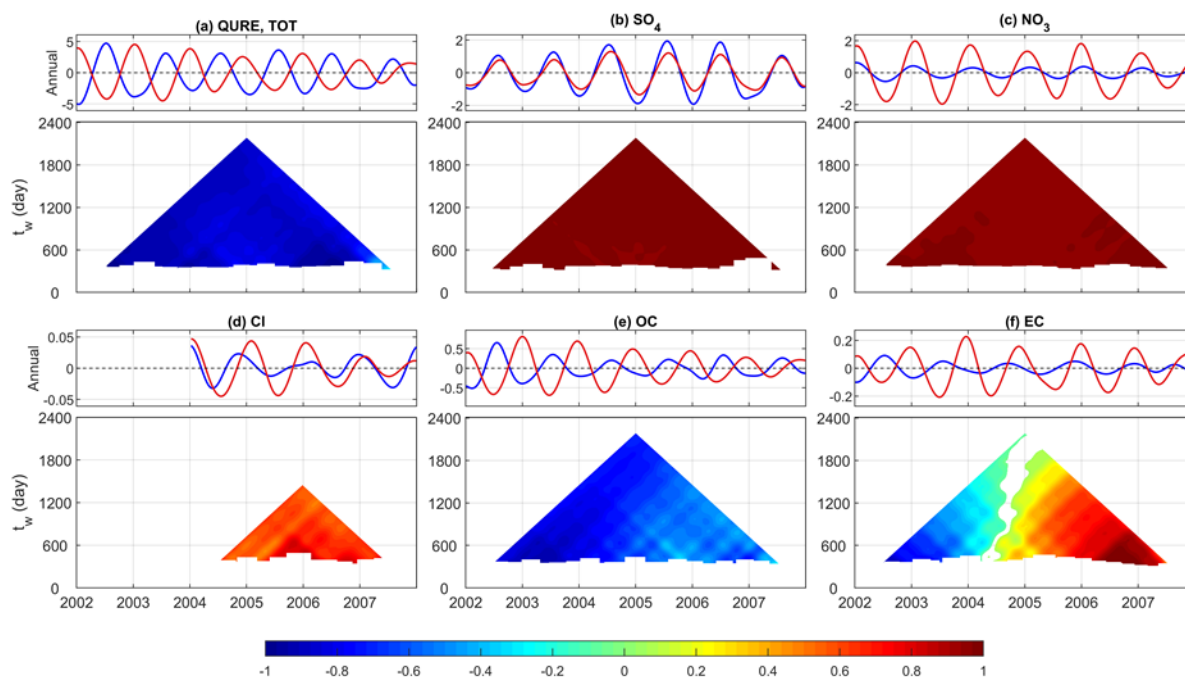


351 During the period of 2002-2007, observations at ATL reveal a slightly increasing  $PM_{2.5}$  trend that cannot be explained  
352 by the five available  $PM_{2.5}$  components trend (Fig. 6c), indicating a contribution of the remaining species such as the  
353 non-carbonaceous portion of organic matter. Non-carbonaceous organic matter can account for more than half of total  
354 organic matter, which, in turn, can account for a large portion of the total  $PM_{2.5}$  mass (Edgerton et al., 2005). In contrast,  
355 the model shows a slight decreasing trend with a peak decreasing rate in 2003 and misses the peak increasing rate of  
356  $0.23 \mu\text{g}/\text{m}^3/\text{year}$  in the winter season of 2005. Similarly, reversed trends are also simulated for  $SO_4$ , OC, and EC, while  
357 the change rate in  $NO_3$  is well captured. Unlike the previous sites, magnitude of trend components in total and  
358 speciated  $PM_{2.5}$  are well simulated except for EC (1.4 times the observation) and  $NO_3$  (2.1 times).

359 To sum up, the decreasing long-term trend at QURE is well simulated by the model. The occurrence of large wildfires  
360 lasting for several months has significantly impacted the long-term trend component at RENO and the model failed  
361 to capture those combustion-related species and total  $PM_{2.5}$  primarily due to limitations in the historical data used to  
362 specify day-specific wildfire emissions (Xing et al., 2013). Slightly increasing levels of  $PM_{2.5}$  and its species observed  
363 at ATL are simulated to be slightly decreasing, except for  $NO_3$  which is well simulated. The magnitude of the long-  
364 term trend components of total  $PM_{2.5}$  and  $SO_4$  are well represented by CMAQ (Table 2). The model performs  
365 differently across the sites in terms of the magnitudes of the trend component in  $NO_3$ ,  $NH_4$ , Cl, OC, and EC. The large  
366 discrepancy in the magnitude of some long-term trend components is likely pointing to the systematic bias in the  
367 annual emission estimations as discussed in Xing et al., (2013), which mainly focused on long-term trend rather than  
368 the absolute level of the emissions. Species other than those in the available dataset also play a considerable role in  
369 driving the agreements or disagreements between model simulations and observations of total  $PM_{2.5}$ .

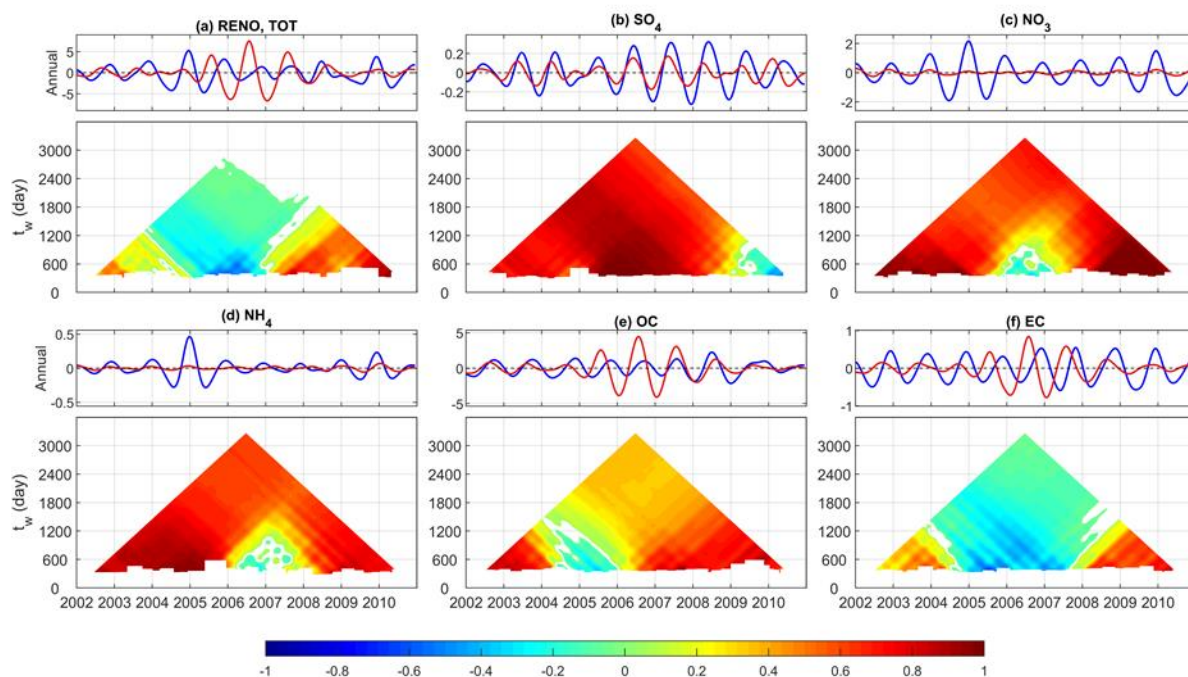
### 370 **4.3 Seasonality**

371 The EMD-assisted seasonality evaluations utilize the decomposed IMFs with characteristic period of one year to  
372 evaluate the amplitude and phase of the model simulation, both of which are time-dependent. As mentioned in Section  
373 4.1, these IMFs are statistically significant from white noise with few exceptions (Fig. S5). We first demonstrate the  
374 evaluation for total  $PM_{2.5}$  at QURE (Fig. 7a). The top panel shows the annual cycle components and the bottom panel  
375 shows its TDIC pyramid. The decreasing amplitude of the annual cycles throughout 2002-2007 is almost perfectly  
376 represented with an overall ratio  $r_{annual}$  being 1.0 (Table 3). Each pixel in the TDIC pyramid is the correlation (color-  
377 coded) calculated during a period of time  $I(t)$  with width of  $t_w$  days (y-axis) centered at a specific day (x-axis) as  
378 introduced in Section 3.2. The annual cycle mean periods are identical between CMAQ and observations (350 days,  
379 Fig. 2a IMF6), but there is a phase shift for all years with the entire TDIC pyramid being close to -1. By shifting the  
380 CMAQ annual cycles backward 159 days (almost half year), the overall correlation of the annual component can reach  
381 up to a peak of 0.9 (Table 4).



382  
 383 **Fig. 7. Decomposed annual cycles (IMF6) from observed (blue) and simulated (red) concentrations ( $\mu\text{g}/\text{m}^3$ ) of**  
 384 **(a) total  $\text{PM}_{2.5}$ , (b)  $\text{SO}_4$ , (c)  $\text{NO}_3$ , (d)  $\text{Cl}$ , (e)  $\text{OC}$ , and (f)  $\text{EC}$  and their corresponding TDIC at Quabbin Summit,**  
 385 **MA. The window size  $t_w$  indicates the width of the window used to calculate a specific correlation centered at**  
 386 **the day represented in x-axis.**

387 What are the driving factors for the above phase shift in modeled total  $\text{PM}_{2.5}$  at Quabbin Summit, MA? The illustrations  
 388 in Fig. 7a for total  $\text{PM}_{2.5}$  alone cannot provide useful information that will allow the modeler to improve the model's  
 389 performance. This is accomplished by applying the EMD method to the  $\text{PM}_{2.5}$  speciated components (Fig. 7b-f). Traces  
 390 of the semi-annual phase shift (-159 days) of annual cycles or large overestimation in the winter and underestimation  
 391 in the summer is because of the largely overestimated amplitude of  $\text{NO}_3$  (4.3 times that of observation) which peaks  
 392 in the winter and the almost semi-annual shifted  $\text{OC}$  (-147 days), as well as contributions from  $\text{EC}$  and  $\text{Cl}$ .  $\text{NO}_3$  has a  
 393 mean amplitude reaching almost half of that of the total  $\text{PM}_{2.5}$ .  $\text{OC}$  directly drives both the observed and simulated  
 394 annual components to be negatively correlated.  $\text{EC}$  follows the feature of  $\text{OC}$  in the first four years or so and the  
 395 feature of  $\text{NO}_3$  in 2006 and 2007 and contributes to the half year shifted total  $\text{PM}_{2.5}$ . The magnitude of winter-peaking  
 396  $\text{Cl}$  cycles is overestimated with a phase shift of one month. However, the contribution of  $\text{Cl}$  is very limited because of  
 397 the tiny amplitude in both observed and simulated annual cycles. In addition, annual cycles in  $\text{SO}_4$  are well reproduced  
 398 for the entire time span with an amplitude ratio of 0.7. A quantitative summary of the evaluation of the annual cycles  
 399 at this site can be found in Tables 3 and 4.



400

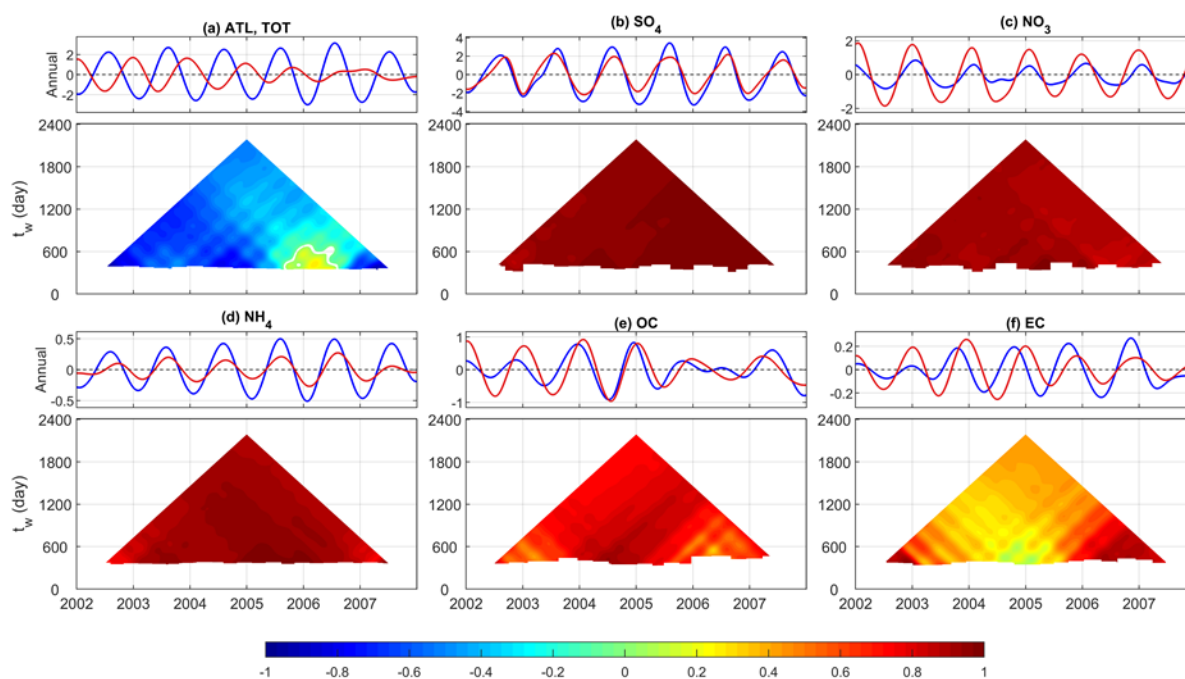
401 **Fig. 8. Same as in Fig. 7 for Reno, NV, except that (d) represents NH<sub>4</sub> rather than Cl.**

402 Both observed and simulated annual cycles at the RENO site are largely influenced by the extreme events lasting for  
 403 several months that are not properly simulated, indicating the need for more accurately specified wildfire emissions.  
 404 Overall, annual variations for total and speciated PM<sub>2.5</sub> are largely underestimated except for the total PM<sub>2.5</sub> and  
 405 combustion-driven EC and OC from 2005 to 2007 (Fig. 8). The modeled phase of SO<sub>4</sub>, NO<sub>3</sub>, NH<sub>4</sub>, and OC agrees  
 406 with that of observation with the exception for a length of about two years in each that missed the phasing: 2009-2010  
 407 for SO<sub>4</sub>, summer 2005-summer 2007 for NO<sub>3</sub>, 2006-2007 for NH<sub>4</sub>, and 2004-2005 for OC. It is also notable that the  
 408 TDIC pyramid of EC mimics that of total PM<sub>2.5</sub>, implying the existence of errors in modeled EC in processes such as  
 409 emissions, transport, and deposition that affected the model performance for total PM<sub>2.5</sub>. In comparison, SO<sub>4</sub> and OC  
 410 are relatively well simulated with a mean amplitude ratio of 0.5 and 1.5 and a phase shift of 36 and 33 days,  
 411 respectively.

412 Observed annual cycles of total PM<sub>2.5</sub> at the ATL site features a slightly increasing amplitude of annual variations  
 413 from 2002 to 2006 which then decreased to the original state in 2007 (Fig. 9a). Conversely, model-simulated annual  
 414 cycles became weaker throughout the period, with an overall  $r_{annual}$  of 0.5. As at the QURE site, the simulated annual  
 415 components at the ATL site also show a shift of several months (-132 days). Specifically, traces of these phase shifts  
 416 or large overestimation in the winter and underestimation in the summer can be seen from the more than doubled  
 417 amplitude of NO<sub>3</sub> which peaks in winter and underestimated SO<sub>4</sub> and NH<sub>4</sub> in the warm seasons as well as the -54 days  
 418 shifted EC. The anti-correlated remaining species other than those in the available dataset clearly played a role in  
 419 driving the discrepancies seen in the total PM<sub>2.5</sub> annual cycles (Fig. 10). Specifically, the anti-correlation likely points  
 420 to an inaccurate representation of the seasonal variation of the non-carbonaceous portion of organic matter due to an  
 421 incomplete representation of organic aerosols in the model version analyzed here; newer versions of the CMAQ model

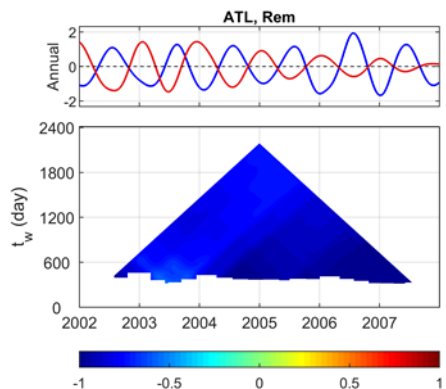
422 include updated treatment of organic aerosols (e.g., additional SOA formation pathways, improvements in  
 423 representation of primary OM emissions) which is likely to improve the mentioned features (Appel et al., 2017;  
 424 Murphy et al., 2017; Xu et al., 2018). The underestimated annual variations in the remaining components closely  
 425 resemble that of the annual variation in total  $PM_{2.5}$ . The phase of simulated  $SO_4$ ,  $NO_3$ ,  $NH_4$ , and OC species is in good  
 426 agreement with those in observations and the amplitude of simulated annual cycles in  $SO_4$ , OC, and EC agree well  
 427 with that in the observations (Tables 3 and 4).

428 In sum, annual cycles of  $PM_{2.5}$  are also time-dependent and the phase in the annual cycles for total  $PM_{2.5}$ , OC, and EC  
 429 reveals a general shift of up to half a year (Table 4); this indicates a potential problem in the allocation of emissions  
 430 during this study period and/or the treatment of organic aerosols in this version of the model. CMAQ generally  
 431 simulated the phase in  $SO_4$ ,  $NO_3$ , Cl, and  $NH_4$  quite well but did not always capture the magnitude of their variations  
 432 (Table 3).



433  
 434 **Fig. 9.** Same as in Fig. 7 for Atlanta, GA, except that the annual component is resolved in IMF8 (IMF7 for  $SO_4$   
 435 and  $NO_3$ ) because of the difference in sampling rate and characteristic embedded in the time series at ATL and  
 436 (d) represents  $NH_4$  rather than Cl.

437



438  
 439 **Fig. 10.** Decomposed annual cycles in Atlanta, GA for the remaining components presented in total PM<sub>2.5</sub> other  
 440 than the five species in Fig.9.

441 **Table 3.** The ratio of mean amplitude of the annual component  $r_{annual}$  (CMAQ/observation). Boldface values indicate  
 442 a magnitude with a ratio close to 1 (0.7-1.3).

	TOT	SO <sub>4</sub>	NO <sub>3</sub>	NH <sub>4</sub>	OC	EC	Cl
QURE	<b>1.0</b>	<b>0.7</b>	4.3	-	1.6	3.1	1.6
RENO	<b>1.2</b>	0.5	0.1	0.2	1.5	<b>0.9</b>	-
ATL	0.5	<b>0.7</b>	2.4	0.4	<b>1.2</b>	<b>1.0</b>	-

443  
 444 **Table 4.** Phase shift ( $n$ ) of CMAQ simulated annual cycle components in days. The background color indicates the  
 445 maximum correlation ( $R_{max}$ ) that can be reached by shifting the CMAQ time series  $n$  days with respect to  
 446 observations: white = [0.8, 1], light grey = [0.6, 0.8), grey = [0.4, 0.6), dark grey = (0.2, 0.4). The bold shows number  
 447 of shifts less than a month while the italic shows shifts longer than three months.

	TOT	SO <sub>4</sub>	NO <sub>3</sub>	NH <sub>4</sub>	OC	EC	Cl
QURE	<i>-159</i>	<b>-6</b>	<b>3</b>	-	<i>-147</i>	<b>-105</b>	<b>-30</b>
RENO	<b>78</b>	36	<b>12</b>	<b>-21</b>	33	96	-
ATL	<i>-132</i>	<b>0</b>	<b>8</b>	<b>-17</b>	<b>-24</b>	-54	-

448  
 449 **4.4 Sub-seasonal and inter-annual variability**

450 In this section, model performance at multiple sub-seasonal and inter-annual scales with cycles less than 3 years,  
 451 presented in the total and speciated PM<sub>2.5</sub>, is evaluated following an approach similar to that for the annual cycles in  
 452 Section 4.3 (Fig. 11). First, IMFs from observations and model simulations are paired based on their characteristic  
 453 periods following the discussion in Section 4.1. Then, the magnitude of specific scales is evaluated using  $r_{IMFn}$   
 454 following Equation 6 of the  $r_{annual}$  for annual cycles. The phase shifts of the time series are assessed by the proportion  
 455 of shifted days relative to the mean characteristic scales of the corresponding observed and simulated IMFs ( $n/t_m$ ).

456 For example, a phase shift of 0.1 cycles in the 2-year cycles is approximately 73 days while it would be 18 days for  
457 the half-year cycles.

458 The performance of the simulated amplitude of the sub-seasonal and inter-annual cycles is relatively stable from a few  
459 days to semi-annual scales and  $r_{IMFn}$  is close to 1 in most cases (Fig. 11a-c). CMAQ captures the features seen in the  
460 observations at QURE, except for the large overestimation of  $\text{NO}_3$  ( $r_{IMFn}$  ranges from 2.6 to 3.7 at the sub-seasonal  
461 scale and reaches up to 13.8 for the 3-year cycles). Similar overestimation of  $\text{NO}_3$  is also found at ATL ( $r_{IMFn}$  ranges  
462 from 2.0 to 3.4, except for the 2-year cycles). In contrast,  $\text{NO}_3$  at RENO is strongly underestimated with  $r_{IMFn}$  ranging  
463 from 0.1 to 0.3 and reaching its minimum at the 2-year cycles. Likewise, all time scales of  $\text{NH}_4$  at RENO are also  
464 being underestimated with  $r_{IMFn}$  decreasing from 0.4 to only 0.1 at the 3-year cycles. The coexistence of  
465 underestimation of  $\text{NO}_3$  and  $\text{NH}_4$  variability, as well as their trend component, likely points to the insufficient grid  
466 resolution in representing ammonium nitrate episodes associated with stagnant meteorology in the mountainous  
467 regions as illustrated by Kelly et al. (2019). To sum up, model has simulated the magnitude of features across all scales  
468 in most of the studied cases. However, fluctuations in  $\text{NO}_3$  are constantly being largely over- or under-estimated and  
469 improvements to the model are required to better replicate its variability (Fig. 11a-c).

470 A high  $R_{max}$  of corresponding IMFs can only be achieved when the characteristic scales of those from observations  
471 and model simulations are close, there is minimal mode mixing, and negligible irregular change of amplitude exists  
472 during the study period. Thus,  $R_{max}$  tends to be small for all oscillations at RENO because of the irregular impact  
473 from events such as wildfires. Thus, the interpretation of phase shift is focused on the components and time scales  
474 having correlations above 0.4 only.

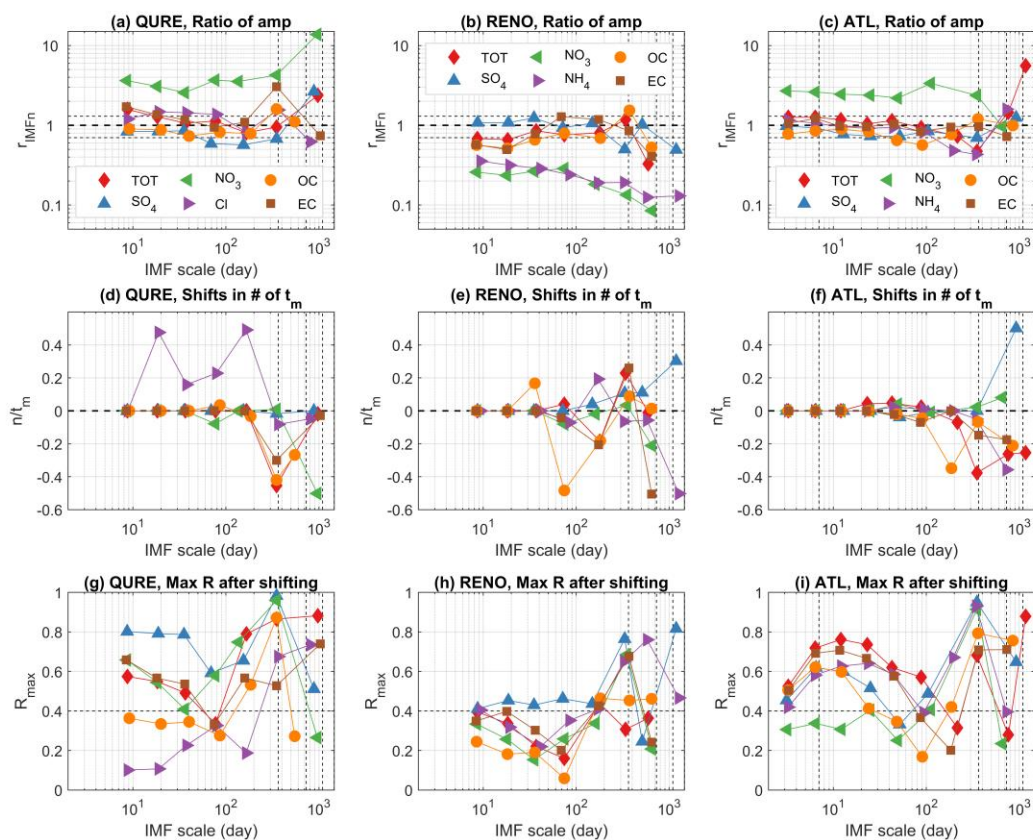
475 Results show that the sub-seasonal cycles at QURE all have a negligible phase shift of less than 0.1 cycles (Fig. 11d).  
476 The semi-annual cycles at RENO have around 0.2 cycle phase shifts in total  $\text{PM}_{2.5}$  (-0.2),  $\text{NH}_4$ (0.2), OC (-0.2), and  
477 EC (-0.2) while negligible phase shifts of less than 0.1 cycles are simulated in  $\text{SO}_4$  ranging from 9 days to semi-annual  
478 in scale. As at QURE, multiple sub-seasonal cycles at ATL all have a negligible phase shift of less than 0.1 cycles,  
479 with the exception of semi-annual OC which has a phase shift of nearly -0.4 cycles with a marginal correlation of  
480 around 0.4. Unlike the relatively stable  $R_{max}$  throughout the time scales within each of the species for QURE and  
481 RENO, the  $R_{max}$  at ATL tends to be much higher (roughly 0.6-0.8) in the scales of 6 to 25 days, except for  $\text{NO}_3$ ,  
482 indicating the model's success in simulating those weather-induced air quality fluctuations at this site as reflected by  
483 their negligible phase shifts.

484 However, the physical meaning of each sub-seasonal IMF is not yet fully understood and requires further study. For  
485 example, synoptic scale IMFs (IMFs with scale less than/around a month) usually have large variance and are not  
486 statistically significantly different from white noise except for observed  $\text{SO}_4$  and  $\text{NH}_4$  (Fig. S5). Yet, observed and  
487 simulated total and some speciated  $\text{PM}_{2.5}$  at QURE and ATL (except IMF1) can achieve moderate to high  $R_{max}$  at  
488 these time scales (Fig. 11 g-i), indicating a potential physical explanation of those time scales using meteorological  
489 variables. IMFs with scales longer than a month but less than half year possess much less variance and are usually not  
490 statistically significantly different from noise. Exceptions are also found at the Atlanta site where observed IMFs are



491 mostly significantly different from noise. Whereas semi-annual cycles are mostly statistically significant (note that  
 492 semi-annual  $\text{SO}_4$  and  $\text{NO}_3$  at ATL are too weak to be decomposed into a separate IMF). In a previous study, He et al.  
 493 (2014) found semi-annual oscillations in the corrected AERosol RObotic NETwork (AERONET) Aerosol Optical  
 494 Depth (AOD) and  $\text{PM}_{10}$  mass concentrations are primarily caused by the change of wind directions in Hong Kong.

495



496

497 **Fig. 11. Model performance at all temporal scales for sites QURE, RENO, and ATL. (a-c) ratio of mean**  
 498 **amplitude of corresponding IMFs with similar characteristic mean periods (ideal ratio=1.0); (d-f) the phase**  
 499 **shift  $n$  in the number of mean periods (average mean period of corresponding IMFs decomposed from**  
 500 **observation and model simulation); (g-i) maximum correlation  $R_{max}$  can be achieved by shifting the modeled**  
 501 **time series. The average mean period of corresponding IMFs decomposed from observations and CMAQ of**  
 502 **total and speciated  $\text{PM}_{2.5}$  are represented on the x-axis; all metrics on the y-axis are unitless. Horizontal**  
 503 **reference lines are drawn at 0.7 and 1.3 in (a-c). Weekly, annual, and inter-annual (2- to 3-year) scales are**  
 504 **marked with vertical dashed lines.**

505 The evaluation and interpretation of inter-annual cycles are constrained by the limited available speciated observations  
 506 for a period of 6 to 9 years (4 years for Cl at QURE). Thus, only 2- to 3-year cycles are presented (Fig. 11) and  
 507 evaluated. Among the 2- to 3-year inter-annual cycles at QURE, there is minimal phase shift for total  $\text{PM}_{2.5}$ ,  $\text{SO}_4$ , Cl,  
 508 and EC with moderate to high  $R_{max}$ . At RENO, the model presents negligible shifts in 2-year cycles of OC and  $\text{NH}_4$

509 while phase shifts of 0.3 and -0.5 cycles are simulated in the 3-year cycles for SO<sub>4</sub> and NH<sub>4</sub>. At ATL, the phase shift  
510 of -0.2 to -0.4 cycles are simulated for PM<sub>2.5</sub>, NH<sub>4</sub>, OC, and EC with periods of 2- to 3-year cycles; while 2- to 3-year  
511 SO<sub>4</sub> cycles have a half-year cycle shift.

## 512 **5 Conclusions**

513 The main advantage for using EMD to evaluate PM<sub>2.5</sub> and its speciated components is that it decomposes nonlinear  
514 and nonstationary signals into multiple modes and a residual trend component. It does not require any preselection of  
515 the temporal scales and assumptions of linearity and stationarity for the data, thereby providing insights into time  
516 series of PM<sub>2.5</sub> concentrations and its components. Using improved CEEMDAN, we are able to assess how well  
517 regional-scale air quality models like CMAQ can simulate the intrinsic time-dependent long-term trend and cyclic  
518 variations in daily average PM<sub>2.5</sub> and its species. This type of coordinated decomposition and evaluation of total and  
519 speciated PM<sub>2.5</sub> provides a unique opportunity for modelers to assess influences of each PM<sub>2.5</sub> species to the total  
520 PM<sub>2.5</sub> concentration in terms of time shifts for various temporal cycles and the magnitude of each component including  
521 the trend.

522 A demonstration of how improved CEEMDAN could be applied to PM<sub>2.5</sub> time series at three sites over CONUS that  
523 provide speciated PM<sub>2.5</sub> data reveals the presence of the annual cycles in PM<sub>2.5</sub> concentrations and time-dependent  
524 features in all decomposed components. At these three sites, the model generally is more capable of simulating the  
525 change rate in the trend component than the absolute magnitude of the long-term trend component. However, the  
526 magnitude of SO<sub>4</sub> trend components is well represented across all three sites. Also, the model reproduced the amplitude  
527 of the annual cycles for total PM<sub>2.5</sub>, SO<sub>4</sub>, and OC. The phase difference in the annual cycles for total PM<sub>2.5</sub>, OC, and  
528 EC reveal a shift of up to half year, indicating the need for proper allocation of emissions and an updated treatment of  
529 organic aerosols compared to the earlier model version used in this set of model simulations. The consistent large  
530 under/over-prediction of NO<sub>3</sub> variability at all temporal scales and magnitude in the trend component, as well as the  
531 abnormally low correlations of synoptic scale NO<sub>3</sub> at ATL, calls for better representation of nitrate partitioning and  
532 chemistry. Wildfires have the potential to elevate PM<sub>2.5</sub> for months and can alter its variability at scales from few days  
533 to the entire year. Thus, more accurate fire emission data should be incorporated to improve model simulation,  
534 especially in those fire-prone regions.

535 **Data availability.** Paired observations and CMAQ model data used in the analysis will be made available at  
536 <https://edg.epa.gov/metadata/catalog/main/home.page>. Raw CMAQ model outputs are available on request from the  
537 U.S. EPA authors.

538 **Author contribution.** "HL and MA designed the methodology; RM, CH, and SR contributed in the assessment of the  
539 outcomes and were consulted on necessary revisions. Model simulations were performed by the US EPA. HL prepared  
540 the manuscript with contributions from all co-authors."

## 541 **Acknowledgements**

542 The views expressed in this paper are those of the authors and do not necessarily represent the view or policies of the  
543 U.S. Environmental Protection Agency. Two of the authors (MA and HL) acknowledge that part of this work was  
544 supported by the Electric Power Research Institute (EPRI) Contract #00-10005071, 2015–2017.

## 545 **References**

546 Appel, K.W., Napelenok, S.L., Foley, K.M., Pye, H.O., Hogrefe, C., Luecken, D.J., Bash, J.O., Roselle, S.J., Pleim,  
547 J.E., Foroutan, H. and Hutzell, W.T., 2017. Description and evaluation of the Community Multiscale Air Quality  
548 (CMAQ) modeling system version 5.1. *Geoscientific Model Development*, 10(4), p.1703.

549 Astitha, M., Luo, H., Rao, S.T., Hogrefe, C., Mathur, R., Kumar, N., 2017. Dynamic evaluation of two decades of  
550 WRF-CMAQ ozone simulations over the contiguous United States. *Atmospheric Environment* 164, 102–116.

551 Banzhaf, S., Schaap, M., Kraneburg, R., Manders, A.M.M., Segers, A.J., Visschedijk, A.H.J., Denier van der on,  
552 H.A.C., Kuenen, J.P.P., van Meijgaard, E., van Ulft, L.H., Cofala, J., Builtjes, P.J.H., 2015. Dynamic model evaluation  
553 for secondary inorganic aerosol and its precursors over Europe between 1990 and 2009. *Geoscientific Model*  
554 *Development* 8, 1047–1070.

555 Chang, P.C., Flatau, A., Liu, S.C., 2003. Review Paper: Health Monitoring of Civil Infrastructure. *Structural Health*  
556 *Monitoring* 2, 257–267.

557 Chen, X., Wu, Z., Huang, N.E., 2010. The time-dependent intrinsic correlation based on the empirical mode  
558 decomposition. *Adv. Adapt. Data Anal.* 02, 233–265.

559 Civerolo, K., Hogrefe, C., Zalewsky, E., Hao, W., Sistla, G., Lynn, B., Rosenzweig, C., Kinney, P.L., 2010. Evaluation  
560 of an 18-year CMAQ simulation: Seasonal variations and long-term temporal changes in sulfate and nitrate.  
561 *Atmospheric Environment* 44, 3745–3752.

562 Colominas, M.A., Schlotthauer, G., Torres, M.E., 2014. Improved complete ensemble EMD: A suitable tool for  
563 biomedical signal processing. *Biomedical Signal Processing and Control* 14, 19–29.

564 Derot, J., Schmitt, F.G., Gentilhomme, V., Morin, P., 2016. Correlation between long-term marine temperature time  
565 series from the eastern and western English Channel: Scaling analysis using empirical mode decomposition. *Comptes*  
566 *Rendus Geoscience* 348, 343–349.

567 Edgerton, E.S., Hartsell, B.E., Saylor, R.D., Jansen, J.J., Hansen, D.A., Hidy, G.M., 2005. The Southeastern Aerosol  
568 Research and Characterization Study: Part II. Filter-Based Measurements of Fine and Coarse Particulate Matter Mass  
569 and Composition. *Journal of the Air & Waste Management Association* 55, 1527–1542.

570 Foley, K.M., Hogrefe, C., Pouliot, G., Possiel, N., Roselle, S.J., Simon, H., Timin, B., 2015. Dynamic evaluation of  
571 CMAQ part I: Separating the effects of changing emissions and changing meteorology on ozone levels between 2002  
572 and 2005 in the eastern US. *Atmospheric Environment* 103, 247–255.

573 Gan, C.-M., Pleim, J., Mathur, R., Hogrefe, C., Long, C.N., Xing, J., Wong, D., Gilliam, R., Wei, C., 2015. Assessment  
574 of long-term WRF–CMAQ simulations for understanding direct aerosol effects on radiation “brightening” in the  
575 United States. *Atmospheric Chemistry and Physics* 15, 12193–12209.

576 Gyawali, M., Arnott, W.P., Lewis, K. and Moosmüller, H., 2009. In situ aerosol optics in Reno, NV, USA during and  
577 after the summer 2008 California wildfires and the influence of absorbing and non-absorbing organic coatings on  
578 spectral light absorption. *Atmospheric Chemistry & Physics*, 9(20).

579 Hansen, D.A., Edgerton, E.S., Hartsell, B.E., Jansen, J.J., Kandasamy, N., Hidy, G.M., Blanchard, C.L., 2003. The  
580 Southeastern Aerosol Research and Characterization Study: Part 1—Overview. *Journal of the Air & Waste*  
581 *Management Association* 53, 1460–1471.

582 He, J., Zhang, M., Chen, X., & Wang, M., 2014. Inter-comparison of seasonal variability and nonlinear trend between  
583 AERONET aerosol optical depth and PM10 mass concentrations in Hong Kong. *Science China Earth*  
584 *Sciences*, 57(11), 2606-2615.

585 Henneman, L.R.F., Liu, C., Hu, Y., Mulholland, J.A., Russell, A.G., 2017. Air quality modeling for accountability  
586 research: Operational, dynamic, and diagnostic evaluation. *Atmospheric Environment* 166, 551–565.

587 Hogrefe, C., Hao, W., Zalewsky, E.E., Ku, J.-Y., Lynn, B., Rosenzweig, C., Schultz, M.G., Rast, S., Newchurch, M.J.,  
588 Wang, L., Kinney, P.L., Sistla, G., 2011. An analysis of long-term regional-scale ozone simulations over the  
589 Northeastern United States: variability and trends. *Atmospheric Chemistry and Physics* 11, 567–582.

590 Huang, N.E., Shen Zheng, Long Steven R., Wu Manli C., Shih Hsing H., Zheng Quanan, Yen Nai-Chyuan, Tung Chi  
591 Chao, Liu Henry H., 1998. The empirical mode decomposition and the Hilbert spectrum for nonlinear and non-  
592 stationary time series analysis. *Proceedings of the Royal Society of London. Series A: Mathematical, Physical and*  
593 *Engineering Sciences* 454, 903–995.

594 Huang, Y., Schmitt, F.G., 2014. Time dependent intrinsic correlation analysis of temperature and dissolved oxygen  
595 time series using empirical mode decomposition. *Journal of Marine Systems* 130, 90–100.

596 Kang, D., Hogrefe, C., Foley, K.L., Napelenok, S.L., Mathur, R., Trivikrama Rao, S., 2013. Application of the  
597 Kolmogorov–Zurbenko filter and the decoupled direct 3D method for the dynamic evaluation of a regional air quality  
598 model. *Atmospheric Environment* 80, 58–69.

599 Kelly, J.T., Koplitz, S.N., Baker, K.R., Holder, A.L., Pye, H.O.T., Murphy, B.N., Bash, J.O., Henderson, B.H., Possiel,  
600 N.C., Simon, H., Eyth, A.M., Jang, C., Phillips, S., Timin, B., 2019. Assessing PM2.5 model performance for the  
601 conterminous U.S. with comparison to model performance statistics from 2007-2015. *Atmospheric Environment* 214,  
602 116872.

603 Mathur, R., Xing, J., Gilliam, R., Sarwar, G., Hogrefe, C., Pleim, J., Pouliot, G., Roselle, S., Spero, T.L., Wong, D.C.,  
604 Young, J., 2017. Extending the Community Multiscale Air Quality (CMAQ) Modeling System to Hemispheric Scales:  
605 Overview of Process Considerations and Initial Applications. *Atmos Chem Phys* 17, 12449–12474.

606 Moghtaderi, A., Borgnat, P., Flandrin, P., 2012. Gap-filling by the empirical mode decomposition, in: 2012 IEEE  
607 International Conference on Acoustics, Speech and Signal Processing (ICASSP). Presented at the 2012 IEEE  
608 International Conference on Acoustics, Speech and Signal Processing (ICASSP), pp. 3821–3824.

609 Murphy, B.N., Woody, M.C., Jimenez, J.L., Carlton, A.M.G., Hayes, P.L., Liu, S., Ng, N.L., Russell, L.M., Setyan,  
610 A., Xu, L. and Young, J., 2017. Semivolatile POA and parameterized total combustion SOA in CMAQv5. 2: impacts  
611 on source strength and partitioning. *Atmospheric Chemistry and Physics*, 17, p.11107.

612 Rato, R.T., Ortigueira, M.D., Batista, A.G., 2008. On the HHT, its problems, and some solutions. *Mechanical Systems  
613 and Signal Processing*, Special Issue: *Mechatronics* 22, 1374–1394.

614 Torres, M.E., Colominas, M.A., Schlotthauer, G., Flandrin, P., 2011. A complete ensemble empirical mode  
615 decomposition with adaptive noise, in: 2011 IEEE International Conference on Acoustics, Speech and Signal  
616 Processing (ICASSP). Presented at the 2011 IEEE International Conference on Acoustics, Speech and Signal  
617 Processing (ICASSP), pp. 4144–4147.

618 White, W.H., 2008. Chemical markers for sea salt in IMPROVE aerosol data. *Atmospheric Environment* 42, 261–  
619 274.

620 Wong, D.C., Pleim, J., Mathur, R., Binkowski, F., Otte, T., Gilliam, R., Pouliot, G., Xiu, A., Young, J.O., Kang, D.,  
621 2012. WRF-CMAQ two-way coupled system with aerosol feedback: software development and preliminary results.  
622 *Geoscientific Model Development* 5, 299–312.

623 Wu, Z., Huang, N.E., 2004. A study of the characteristics of white noise using the empirical mode decomposition  
624 method. *Proceedings of the Royal Society of London. Series A: Mathematical, Physical and Engineering Sciences*  
625 460, 1597–1611.

626 Wu, Z., Huang, N.E., 2009. Ensemble empirical mode decomposition: a noise-assisted data analysis method. *Adv.  
627 Adapt. Data Anal.* 01, 1–41.

628 Wu, Z., Huang, N.E., Long, S.R., Peng, C.-K., 2007. On the trend, detrending, and variability of nonlinear and  
629 nonstationary time series. *PNAS* 104, 14889–14894.

630 Xing, J., Mathur, R., Pleim, J., Hogrefe, C., Gan, C.-M., Wong, D.C., Wei, C., Gilliam, R., Pouliot, G., 2015.  
631 Observations and modeling of air quality trends over 1990–2010 across the Northern Hemisphere: China, the United  
632 States and Europe. *Atmospheric Chemistry and Physics* 15, 2723–2747.

633 Xing, J., Pleim, J., Mathur, R., Pouliot, G., Hogrefe, C., Gan, C.-M., Wei, C., 2013. Historical gaseous and primary  
634 aerosol emissions in the United States from 1990 to 2010. *Atmospheric Chemistry and Physics* 13, 7531–7549.

635 Xu, L., Pye, H.O., He, J., Chen, Y., Murphy, B.N. and Ng, L.N., 2018. Experimental and model estimates of the  
636 contributions from biogenic monoterpenes and sesquiterpenes to secondary organic aerosol in the southeastern United  
637 States. *Atmospheric chemistry and physics*, 18(17), p.12613.

- 638 Yahya, K., Wang, K., Campbell, P., Glotfelty, T., He, J., Zhang, Y., 2016. Decadal evaluation of regional climate, air  
639 quality, and their interactions over the continental US using WRF/Chem version 3.6.1. *Geoscientific Model*  
640 *Development* 9, 671–695.
- 641 Yeh, J.-R., Shieh, J.-S., Huang, N.E., 2010. Complementary ensemble empirical mode decomposition: a novel noise  
642 enhanced data analysis method. *Adv. Adapt. Data Anal.* 02, 135–156.
- 643 Yu, L., Wang, S., Lai, K.K., 2008. Forecasting crude oil price with an EMD-based neural network ensemble learning  
644 paradigm. *Energy Economics* 30, 2623–2635.
- 645 Zhou, W., Cohan, D.S., Napelenok, S.L., 2013. Reconciling NO<sub>x</sub> emissions reductions and ozone trends in the U.S.,  
646 2002–2006. *Atmospheric Environment* 70, 236–244.
- 647