# Evaluating Trends and Seasonality in Modeled PM<sub>2.5</sub> Concentrations Using Empirical Mode Decomposition

Huiying Luo<sup>1</sup>, Marina Astitha<sup>1\*</sup>, Christian Hogrefe<sup>2</sup>, Rohit Mathur<sup>2</sup>, S. Trivikrama Rao<sup>1,3</sup> 3 4 <sup>1</sup>University of Connecticut, Department of Civil and Environmental Engineering, Storrs-Mansfield, CT, USA 5 <sup>2</sup>U.S. Environmental Protection Agency, Research Triangle Park, NC, USA 6 <sup>3</sup>North Carolina State University, Raleigh, NC, USA 7 \*Corresponding author: Marina Astitha, Civil and Environmental Engineering, University of Connecticut, 261 Glenbrook Road, Storrs, CT, 06269-3037, Phone: 860-486-3941, Fax: 860-486-2298, Email: 8 9 marina.astitha@uconn.edu. 10 Abstract. Regional-scale air quality models are being used for studying the sources, composition, transport,

transformation, and deposition of fine particulate matter (PM<sub>2.5</sub>). The availability of decadal air quality simulations provides a unique opportunity to explore sophisticated model evaluation techniques rather than relying solely on traditional operational evaluations. In this study, we propose a new approach for process-based model evaluation of speciated PM<sub>2.5</sub> using improved Complete Ensemble Empirical Mode Decomposition with Adaptive Noise (improved CEEMDAN) to assess how well version 5.0.2 of the coupled Weather Research and Forecasting model - Community Multiscale Air Quality model (WRF-CMAQ) simulates the time-dependent long-term trend and cyclical variations in the daily average PM<sub>2.5</sub> and its species, including sulfate (SO<sub>4</sub>), nitrate (NO<sub>3</sub>), ammonium (NH<sub>4</sub>), chloride (Cl), organic

18 carbon (OC) and elemental carbon (EC). The utility of the proposed approach for model evaluation is demonstrated

using  $PM_{2.5}$  data at three monitoring locations. At these locations, the model is generally more capable of simulating

20 the rate of change in the long-term trend component than its absolute magnitude. Amplitudes of the sub-seasonal and

21 annual cycles of total PM<sub>2.5</sub>, SO<sub>4</sub> and OC are well reproduced. However, the time-dependent phase difference in the

- 22 annual cycles for total PM<sub>2.5</sub>, OC and EC reveal a phase shift of up to half year, indicating the need for proper temporal
- 23 allocation of emissions and for updating the treatment of organic aerosols compared to the model version used for this
- 24 set of simulations. Evaluation of sub-seasonal and inter-annual variations indicates that CMAQ is more capable of
- 25 replicating the sub-seasonal cycles than inter-annual variations in magnitude and phase.

### 26 Keywords

- 27 Model evaluation, coupled WRF-CMAQ, improved Complete Ensemble Empirical Mode Decomposition (EMD)
- 28 with Adaptive Noise, Speciated PM<sub>2.5</sub>, Scale Separation, Seasonality, Trend
- 29

#### 30 **1 Introduction**

31 It is well recognized that inhalable fine particulate matter  $(PM_{2.5})$  adversely impacts human health and the

environment. Regional-scale air quality models are being used in health impact studies and decision-making related
 to PM<sub>2.5</sub>. Long-term model simulations of PM<sub>2.5</sub> concentrations using regional air quality models are essential to

to  $PM_{2.5}$ . Long-term model simulations of  $PM_{2.5}$  concentrations using regional air quality models are essential to identify long-term trends and cyclical variations such as annual cycles in areas larger than what is covered by in-situ

35 measurements. However, total  $PM_{2.5}$  concentrations are challenging to predict because of the dependence on the

36 contributions from individual PM<sub>2.5</sub> components, such as sulfates, nitrates, carbonaceous species, and crustal elements.

37 In this context, a detailed process-based evaluation of the simulated speciated  $PM_{2.5}$  must be carried out to ensure

38 acceptable replication of observations so model users can have confidence in using regional air quality models for

39 policy-making. Furthermore, process-based information can be useful for making improvements to the model.

Some of the trend or step change evaluations of regional air quality models in the past have focused on specific pairs of years (Kang et al., 2013; Zhou et al., 2013; Foley et al., 2015). These studies do not properly account for the sub-seasonal and inter-annual variations between those specific periods. Trend evaluation is commonly done by linear regression of indexes such as the annual mean or specific percentiles, assuming linearity and stationarity of time series (Civerolo et al., 2010; Hogrefe et al., 2011; Banzhaf et al., 2015; Astitha et al., 2017). The problem with the linear trend evaluation is that there is no guarantee the trend is actually linear during the period of the study because the

46 underlying processes are in fact nonlinear and nonstationary (Wu et al., 2007).

47 Seasonal variations are usually studied and evaluated by investigating the monthly or seasonal means of total and/or 48 speciated PM<sub>2.5</sub> (Civerolo et al., 2010; Banzhaf et al., 2015; Yahya et al., 2016; Henneman et al., 2017). Evaluation of 49 ten-year averaged monthly mean (i.e., ten-year averaged mean in Jan., ..., Dec.) of PM<sub>2.5</sub> simulated with WRF/Chem 50 against the Interagency Monitoring of Protected Visual Environments (IMPROVE) by Yahya et al. (2016) shows that 51 the model captures the observed features of summer peaks in  $PM_{2.5}$  with a phase shift of few months. However, 52 according to the analysis (Fig. 10) in Henneman et al. (2017), the seasonality shown in monthly-averaged PM<sub>2.5</sub> time series is much less distinguishable compared with that of ozone and CMAQ (version 5.0.2) does not replicate the 53 54 monthly PM<sub>2.5</sub> quite well with large underestimation in the summer months. In these studies, the seasonality might not 55 be well represented by the preselected averaging window size of one or three months. In addition, averaging of those 56 monthly or seasonal means across multiple years may conceal the long-term trends or interannual variations driven 57 by climate change, emission control policies or other slow varying processes.

58 To address the above-mentioned problems, we propose a new method for conducting air quality model evaluation for 59 PM<sub>2.5</sub> using improved CEEMDAN. Improved CEEMDAN is an Empirical Mode Decomposition (EMD)-based, data-60 driven intrinsic mode decomposition technique that can adaptively and recursively decompose a nonlinear and 61 nonstationary signal into multiple modes called intrinsic mode functions (IMFs) and a residual (trend component) 62 (Huang et al., 1998; Wu and Huang, 2009; Yeh et al., 2010; Torres et al., 2011; Colominas et al., 2014). It does not 63 require any preselection of the temporal scales or assumptions of linearity and stationarity for the data, thereby 64 providing some insights into time series of PM<sub>2.5</sub> concentrations and its components. Decomposed PM<sub>2.5</sub> long-term trend components and annual cycles from observed and simulated PM2.5 serve as the intuitive carrier of the trend and 65

seasonality evaluation. In the meantime, several other IMFs with characteristic time scales ranging from multiple days

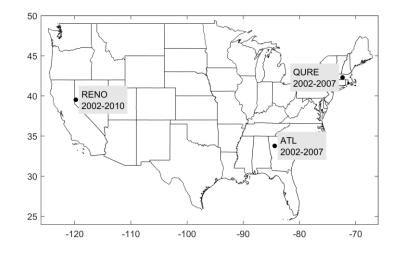
67 to years are also decomposed, enabling model evaluation of the less studied sub-seasonal and inter-annual variations.

#### 68 2 Coupled WRF-CMAQ PM<sub>2.5</sub> Simulations and Observations

69 The two-way coupled WRF-CMAO (version 5.0.2) is configured with a 36 km horizontal grid spacing over the 70 contiguous United States (CONUS) with 35 vertical layers of varying thickness extending from the surface to 50 mb 71 (Wong et al., 2012; Gan et al., 2015). Time-varying chemical lateral boundary conditions were derived from the 108 72 km resolution hemispheric WRF-CMAQ (Mathur et al., 2017) simulation for the 1990-2010 period (Xing et al., 2015). 73 The simulations are driven by a comprehensive emission dataset which includes aerosol precursors and primary 74 particulate matter (Xing et al., 2013, 2015). Annual emissions for the CMAQ simulations were estimated using the 75 methodology described in Xing et al. (2013). Briefly, the National Emissions Inventory (NEI) for 1990, 1995, 1996, 76 1999, 2001, 2002 and 2005 and a number of sector-specific long-term databases containing information about trends 77 in activity data and emission controls were used to create county-level annual emissions for a total of 49 emission 78 sectors. Prior to being used as input to the CMAQ simulations, these annual emissions were then temporally and 79 spatially allocated to provide hourly emissions based on monthly, weekly, and diurnal temporal cross-reference and 80 profile data from the 2005 NEI modeling platform. These profile data vary by emissions source and sometimes by 81 state and county and are generally based on surveys and extrapolation of activity data which can be subject to 82 uncertainty. Exceptions to the use of 2005 NEI platform temporal profile data for temporal allocation were emissions 83 from electric generating units (EGU) which directly used measured hourly emissions after 1995 and wildfire emissions 84 that used climatological monthly, weekly, and diurnal profiles for temporal allocation. Readers can refer to Gan et al. 85 (2015) for additional model information and the trend evaluation against seven pairs of sites from the CASTNET 86 (Clean Air Status and Trend Network) and IMPROVE networks for 1995-2010. We obtained the 2002-2010 daily 87 average  $PM_{2.5}$  and its speciated time series from the set of simulations with direct aerosol feedback. The earlier years 88 of 1990-2001 are not included in this evaluation because of the limited availability of speciated  $PM_{2.5}$  observations.

89 To avoid misinterpretation of data due to the presence of missing values, only sites with continuous complete long-90 term record for total PM<sub>2.5</sub> and its speciation including SO<sub>4</sub>, NO<sub>3</sub>, NH<sub>4</sub>, OC, EC and Cl are studied (Fig. 1). All of the 91 selected sites have data coverage above 90% each year for at least six consecutive years between 2002 and 2010 92 (equivalent to 30% for 1-in-3 days sampling sites). This strict data selection led to the sparsity of this type of 93 observations for the study period. QURE, a rural site carrying out 1-in-3 days sampling of total and speciated PM2.5 94 of SO<sub>4</sub>, NO<sub>3</sub>, OC, EC and Cl, is located in Quabbin Summit, MA. It is one of the three sites from the IMPROVE 95 network that has at least six continuous years of speciated observations and was selected here to demonstrate the 96 application of the proposed method in rural areas. It should be noted that the majority of the observed Cl in 2002 and 97 2003 is negative due to a filter issue problem which was not addressed until 2004 (White, 2008). Thus, simulations of 98 Cl are only evaluated during 2004-2007 at this site. Station RENO, located in urban Reno, NV, is also a 1-in-3 days 99 sampling site of total and speciated PM<sub>2.5</sub> of SO<sub>4</sub>, NO<sub>3</sub>, NH<sub>4</sub>, OC and EC, and it is the only Chemical Speciation 100 Network (CSN) site that fulfills this data coverage requirement. The third site ATL in the Southeastern Aerosol 101 Research and Characterization Study (SEARCH) network is located 4.2 km northwest of downtown Atlanta, GA. It

- is the only long-term site available with daily sampling rate (Hansen et al., 2003; Edgerton et al., 2005) that meets the
- data coverage requirement. The best-estimate (BE), a calculated concentration intended to represent what is actually
- in the atmosphere (Edgerton et al., 2005), of the total PM<sub>2.5</sub> and SO<sub>4</sub>, NO<sub>3</sub>, NH<sub>4</sub> and EC components are retrieved for
- the evaluation. OC component is a direct measurement. These three sites have a continuous record covering at least 6
- 106 years (2002 2007 for QURE and ATL and 2002 2010 for RENO) that allows an evaluation of long-term trends.



108 Fig. 1. Location and data coverage of the PM<sub>2.5</sub> monitoring sites QURE, RENO and ATL.

#### 109 **3 Methodology**

#### 110 **3.1 Empirical Mode Decomposition**

111 The Empirical Mode Decomposition (EMD) technique, proposed in the late 1990s, is capable of adaptively and 112 recursively decomposing a signal into multiple modes called intrinsic mode functions (IMFs), where each mode has 113 a characteristic frequency, and a residual with at most one extremum (Huang et al., 1998). EMD decomposes the 114 original signal into several IMFs and a residual through a repeated process called "sifting": first, local maxima and 115 minima are identified and interpolated separately with a cubic spline as the upper and lower envelop; then an IMF 116 candidate is derived by subtracting the mean of the envelops from the original signal. If the candidate satisfies the 117 following criteria (Huang et al., 1998), it is saved as the first IMF (IMF1), and the remaining portion (original signal 118 - IMF1) is treated as a new input signal for the decomposition of the remaining IMFs; otherwise, more sifting 119 processes should be carried out until the candidate becomes an IMF.

- 120 1) The number of extrema (maxima and minima) and the number of zero-crossings must be equal or differ at most by
- 121 one; 2) The local mean at any point, the mean of the envelope defined by local maxima and the envelope defined by
- 122 local minima, must be zero.
- 123 In this way, IMF1, IMF2, ... are decomposed recursively with decreasing characteristic frequency. The final remaining
- 124 residual (trend) could be a monotonic function of time or a long-term component with one extremum at most. The
- decomposed signal then is expressed as the summation of all IMFs and the final residual:

$$x = \sum_{i=1}^{k} d_i + r \tag{1}$$

127 where x is the original signal,  $d_i$  is the *i*<sup>th</sup> IMF, k is the total number of IMFs and r is the final residual.

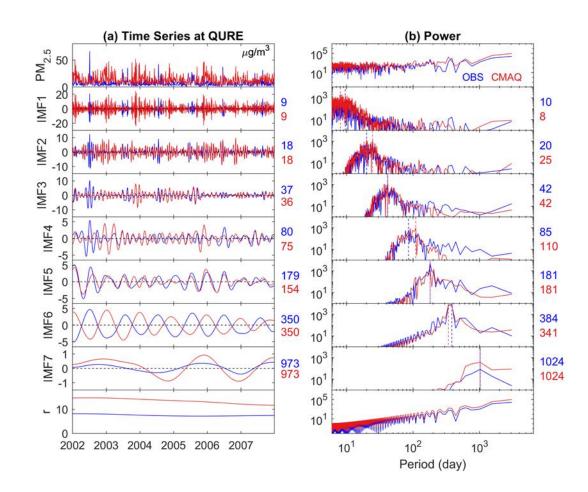
128 Nevertheless, "mode mixing", where oscillations with very disparate scales can be present in one mode or vice versa, 129 is commonly reported. To cope with this issue, multiple noise assisted EMD have been developed successively (Wu 130 and Huang, 2009; Yeh et al., 2010; Torres et al., 2011; Colominas et al., 2014). It is evident that the latest improved 131 Complete Ensemble EMD with Adaptive Noise (improved CEEMDAN) manages to alleviate the problem of mode 132 mixing with the benefit of reducing the amount of noise presented and avoiding spurious modes (Colominas et al., 133 2014). Moreover, the end effects or boundary effects have been addressed by its predecessor EEMD (Ensemble 134 Empirical Mode Decomposition) by extrapolating the maxima and minima, and behaved well in numerous time series 135 with dramatically variant characteristics (Wu and Huang, 2009). The extrapolation of maxima and minima is proven 136 to be more effective compared with the extrapolation of the signal itself such as repetition or reflection (Rato et al., 137 2008).

138 Given the EMD's ability to deal with real-world nonstationary and nonlinear time series data, it is widely used in 139 engineering, economics, earth and environmental sciences (e.g., Huang et al., 1998; Chang et al., 2003; Yu et al., 2008; 140 Colominas et al., 2014; Derot et al., 2016). We use the most up-to-date noise-assisted improved CEEMDAN technique 141 with at least hundreds of noise realizations to decompose observed and simulated  $PM_{2.5}$  time series. Readers can refer 142 to Colominas et al. (2014) for a detailed description of the technique and access to the corresponding MATLAB code. 143 Trial and error attempts are made in setting the inputs (standard deviation of the added noise and the limit of maximum 144 number of sifting allowed) of the improved CEEMDAN function to achieve best mode separation. In a desired best 145 mode separation, neighboring IMFs should have very limited levels of mode mixing, which can be fast screened based 146 on the time series of the decomposed IMFs and their power spectrum.

147 The impact of boundaries on the decomposed annual cycles and the residual is assessed by the variations (standard 148 deviation) of hypothetical decomposed boundaries by cutting a continuous eighteen-year total PM<sub>2.5</sub> observation 149 (North Little Rock, AR) 48 times at different years and times of the year (Fig. S1). The standard deviation is found to 150 largely diminish within half the annual cycles and could be negligible within one year for the annual cycle. This could 151 very possibly expand to IMFs with other characteristic scales. Yet, trend components (residuals) show variability 152 depending on the available time period after cutting. Most of the time, they follow the reference long-term trend 153 reflected either by the residual or the summation of the residual and the IMF with the longest temporal scale 154 decomposed from the eighteen-year  $PM_{2.5}$  (Fig. S1c). This is in line with our expectations as a trend should exist 155 within a given time span, following the definition in Wu et al. (2007): "The trend is an intrinsically fitted monotonic 156 function or a function in which there can be at most one extremum within a given data span". Although very strict 157 data completeness requirement is employed for this study, it should not be conceived as a limitation of the method 158 itself. A sensitivity test based on a period of nine years of total PM<sub>2.5</sub> observations at the same site with 99% data 159 coverage shows that even though variability of annual cycles and long-term trends increases with decreased data 160 availability (100%, 90%..., 10%), the structure of those components is consistent. The average of 40 realizations of 161 annual cycles and long-term trend components in each data-completeness scenario is in perfect alignment with that of

162 100% data completeness (Fig. S2 and S3). Given the fact that those 40 realizations in each scenario are based on 163 independent random samplings of the original observations, the increased variability could very possibly result from 164 the difference in the sampled data itself rather than the method. Thus, the robustness of improved CEEMDAN 165 decomposed annual cycles and long-term trend is justified. In fact, EMD has been proven to be an effective tool for 166 data gap-filling (Moghtaderi et al., 2012).

167



168

169 Fig. 2. Decomposition of observed (blue) and simulated (red) 24-hour average total PM2.5 into 7 IMFs and a 170 residual component (trend) at Quabbin Summit, MA using the improved CEEMDAN: (a) Time series of total 171 PM<sub>2.5</sub>, IMFs and the residual component (all with the unit of  $\mu g/m^3$ ); (b) Power spectrum of the corresponding 172 time series. The colored numbers on the right side of time series are the mean period  $t_m$  in days, while the ones 173 on the right side of the power spectrum are the peak period  $t_p$  in days, which are also indicated by the dashed vertical lines on the power spectrum. Note that the scales for the time series are not all the same. Also, all power 174 175 spectra are in the log scale, and those of the IMFs are zoomed in with a range of  $10^{0}$  to  $10^{4}$  on the y-scale for better visual clarity (compared with  $10^{-2}$  to  $10^{7}$  for total PM<sub>2.5</sub> and the residual component). 176

177

178 The characteristic period of each IMF can be estimated by the peak period  $t_p$  (days) where the power spectrum of the

179 IMF peaks:

$$t_p = \frac{1}{f_p} \qquad (2)$$

in which  $f_p$  is the frequency that the power spectrum peaks in the unit of number of cycles per day. The peak estimates can be biased if more than one high-power frequency is located closely within one IMF. Thus, the power spectrum and  $t_p$  is only used as a fast screening tool to determine if a desired decomposition is accomplished. As an alternative approach, the mean period  $t_m$  can be estimated by:

185 
$$t_m = \frac{\text{Time span}}{(n_{max} + n_{min} + n_{zero})/4}$$
(3)

186 where  $n_{max}$ ,  $n_{min}$  and  $n_{zero}$  are the number of maxima, minima and zero-crossings, respectively, during the 187 *Time span* (days). As the frequency decreases, the mean period estimates become less accurate because of the limited 188 time span compared with the length of the cycle and should be carefully interpreted.

189 An example of the total PM<sub>2.5</sub> decomposition with improved CEEMDAN at the QURE site shows modes ranging from 190 very high frequency to very low frequency (IMF1 to IMF7) and a residual (Fig. 2). No visible mode mixing can be 191 detected in both the time series (Fig. 2a) and the power spectrum (Fig. 2b) of all IMFs. Mean  $(t_m)$  and peak  $(t_p)$ 192 estimations of the characteristic periods of each IMF are presented on the right side of each mode. Annual cycles and 193 long-term trend components are well represented by IMF6 and the residual, with the remaining IMFs carrying weekly, 194 sub-seasonal, seasonal, and inter-annual variations, respectively, for both observed and simulated PM<sub>2.5</sub> (Fig. 2). We 195 have noticed that in some rare cases, a spurious mode in the last IMF with synchronous signal and very close scales 196 to its previous IMF exists. This is possibly due to the fact that the characteristic periods of those IMFs are in proximity 197 to the span of the studied time span. In these cases, the last two modes are merged by adding them together to conduct 198 a detailed evaluation as discussed in Section 4.1.

#### **3.2 Statistical metrics**

200 EMD-decomposed IMFs and trend components allow for a detailed time-dependent evaluation of PM<sub>2.5</sub> and provide 201 a novel opportunity to trace the performances of specific scales back to the corresponding speciated components. Note 202 that the trend component is the decomposed residual component from the  $PM_{2.5}$  in the unit of  $\mu g/m^3$ , and it is not the 203 traditional concept of trend in concentration per time. In addition to a direct evaluation of its magnitude, we also 204 calculated its derivative to identify the periods with higher or lower rate of change (concentration per time). Time-205 dependent intrinsic correlation (TDIC) is utilized to study the evolvement of the model performance for cyclic 206 variations throughout time (Chen et al., 2010; Huang and Schmitt, 2014; Derot et al., 2016). It is a set of correlations 207 calculated for IMFs over a local period of time *I* centered around time *t*:

208 
$$I(t) = [t - \frac{t_w}{2}, t + \frac{t_w}{2}]$$
 (4)

in which t is the center time for the calculation of the correlation and  $t_w$  is the moving window length. The minimum of  $t_w$  is set to be the local instantaneous period of the IMF (larger of that in observation or simulation) using the general zero-crossing method to ensure that at least one instantaneous period is included in calculating the local correlation coefficient (Chen et al., 2010). The maximum of  $t_w$  is the entire data period with a traditional overall correlation being calculated. The empty spaces in the pyramids used to depict the TDIC are an indication that the correlation is not statistically significantly different from zero. With both decomposed observed and modeled concentrations in a narrow scale range, the correlation would no longer be contaminated by coexisting signals of different scales (Chen et al., 2010).

In order to summarize the performance of the decomposed trend component and IMFs, the ratio of the meanmagnitudes of the trend components is defined as:

219 
$$r_{trend} = \frac{Mean_{CMAQ}}{Mean_{observation}}$$
(5)

where *Mean<sub>CMAQ</sub>* and *Mean<sub>observation</sub>* represent the mean of simulated and observed residual components
 respectively. The ratio of the mean amplitude of each IMF is defined by Equation 6, where an example for the annual
 cycles is provided:

223 
$$r_{annual} = \frac{RMS_{CMAQ,annual}}{RMS_{observation,annual}}$$
(6)

where  $RMS_{observation,annual}$  and  $RMS_{CMAQ,annual}$  represent the root mean square of observed and simulated annual cycles respectively. Finally, the phase shift of an IMF *n* is defined as the days an IMF decomposed from modeled time series has to be shifted to maximize the correlation  $(R_{max})$  with the corresponding IMF from observed PM<sub>2.5</sub> time series. In practice, *n* could be as much as a few cycles of the mean period,  $t_m$ . Here, we limit the absolute number of shift days to not exceed a half cycle as a reference for the phase shift of an IMF. Thus, *n* satisfies  $-\binom{t_m}{2} \le n \le$  $\binom{t_m}{2}$  with  $t_m$  being the larger mean period in observation or simulation. It becomes  $-0.5 \le \frac{n}{t_m} \le 0.5$  in terms of number of cycles.

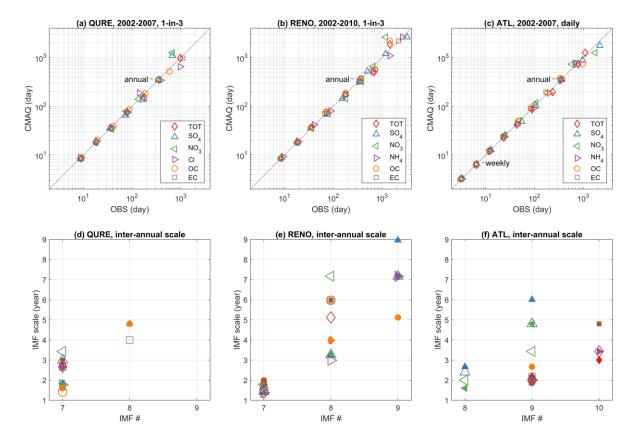
#### 231 4 Results and Discussion

#### 232 4.1 Temporal scales

233 Temporal scales in PM<sub>2.5</sub> resolved by EMD depend solely on the intrinsic properties of the data itself. These properties 234 include underlying characteristics of specific  $PM_{25}$  concentrations, the data sampling frequency, which determines the 235 scales that can be resolved in the high frequency IMFs, and the time span for the data coverage, which could possibly 236 play an important role in differentiating the low frequency IMFs from the trend component. Here, we first evaluate 237 the scales represented by the mean period in the speciated and total PM2.5 time series. Since each IMF represents a 238 nonstationary process, the mean period  $t_m$  is only an estimate of its characteristic scale. Evaluation of  $t_m$  might not 239 necessarily be able to identify issues with corresponding model simulations, and it does not indicate any information 240 on the magnitude or the phase of the time series, which is more important and will be further discussed in Sections 241 4.3 to 4.4.

Fig. 3a presents the characteristic scales  $(t_m)$  of IMFs in observed and simulated total and speciated PM<sub>2.5</sub> of QURE. The CMAQ model compares well with the observations for IMFs 1 through 6 with cycles of 9, 19, 37, 78, 158 and

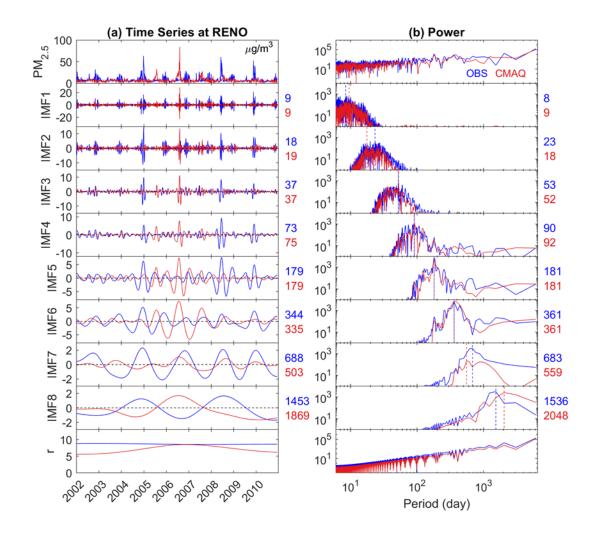
- 244 347 days (average of all observed and simulated total and speciated  $PM_{2.5}$ ). Among all these IMFs, IMF6, which 245 represents the annual cycles, shows the least variations in the characteristic scale (Fig. 3a) and highest peak energy
- from the power spectrum such as Fig. 2b for total PM<sub>2.5</sub>, except for observed EC and OC where the power of half-
- 247 year cycles is more dominant (Fig. S4). These two features demonstrate a clear seasonality in both observed and
- simulated total and speciated  $PM_{2.5}$ , which would otherwise be concealed by practices such as monthly averaging.
- 249 This can be further confirmed by the statistically significant annual cycles (except for observed EC and OC) (Fig. S5)
- based on a Monte Carlo verified relationship between the energy density and mean period of IMFs (Wu and Huang,
- 251 2004; Wu et al., 2007). To explore the inter-annual cycles in more detail, mean periods of IMFs with scales longer
- than a year are being displayed in the top left panel of Fig. 3a. Some variability exists between the observation and
- 253 model simulation to the extent that not all IMFs from observation are being simulated and vice versa for the inter-
- annual cycles. The characteristic scales of all decomposed IMFs with scales longer than a year are shown in Fig. 3d.
- 255 The estimated mean periods of the inter-annual cycles and the differences in the presence of slow varying cycles with
- the long characteristic scales are likely to be influenced by their proximity to the data time span of 6 years (4 years for
- 257 Cl). This implies that the model evaluation shouldn't go beyond 3 years (2 years for Cl) given the current data
- coverage. CMAQ captured the 3-year cycles in EC and total PM<sub>2.5</sub> and 2-year cycles in OC and Cl, despite an
- 259 overestimation in the scales of 2-year cycles in observed SO<sub>4</sub> and NO<sub>3</sub>.
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- 261
- 262
- 263



264

Fig. 3. The characteristic scales  $(t_m)$  resolved in the IMFs of observed and simulated total and speciated PM<sub>2.5</sub> for (a, d) QURE, (b, e) RENO and (c, f) ATL. In (a-c), IMF1 to the last pair of IMFs with increasing characteristic periods are shown from bottom left to top right. Mean periods of IMFs with scales longer than a year are being displayed in (d-f) with the same shapes as in the legend above to show the characteristic scales of all decomposed IMFs given that not all IMFs from observation are being simulated and vice versa. In the (df), species decomposed from observations are shown with smaller filled shapes, while species decomposed from simulations are represented by larger open shapes in slightly darker shades.

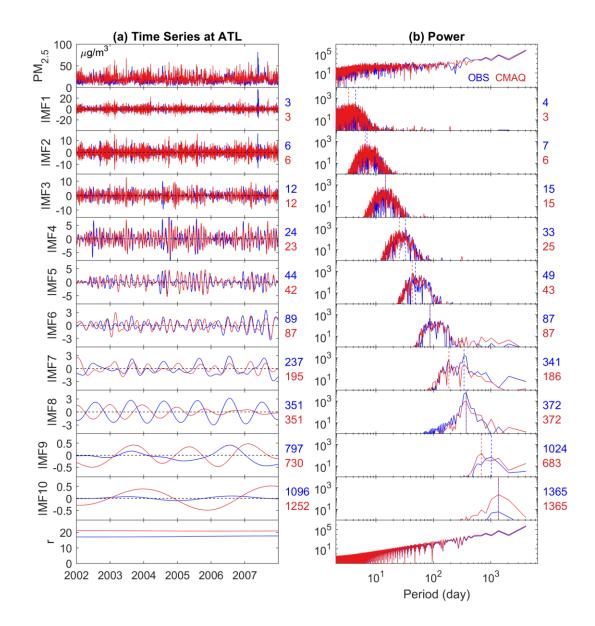
272 Similar features in observed and simulated total and speciated PM<sub>2.5</sub> concentrations at RENO are presented in Fig. 3b. 273 Likewise, the highest peaks in the power spectrum also sit in the annual cycles of IMF6 except for the observed OC 274 and total PM<sub>2.5</sub> which have higher peak power at half-year cycles. All annual IMFs are statistically significant except 275 for simulated NH<sub>4</sub> (Fig. S5). The small variation in the estimated characteristic period of IMF6 is because this 276 monitoring site is located in a wildfire prone region on the border of Nevada and California. Clear evidence can be 277 seen from Fig. 4a that an extra annual cycle in the IMF6 of observations in the summer of 2008 is depicted, which is 278 very possibly driven by the 2008 California Wildfires spanning from May until November. Satellite image of the 279 wildfire smoke on July 10, 2008 can be found in Figure 1 from Gyawali et al. (2009). Unlike the diversified scales in 280 IMF7 at OURE, IMF7 at RENO features universal 2-year cycles of all species as well as total PM25 and all of them 281 are well replicated by the model. However, variations in time scales are present in IMF8 possibly because of the 282 limited data coverage. Thus, only species with time scales less than 4 years in both observations and model simulations 283 are evaluated. It is evident that CMAQ has reproduced the 3-year cycles in SO<sub>4</sub> and NH<sub>4</sub>.



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Fig. 4. Same as Fig. 2 but for the RENO site with 8 IMFs.

287 ATL is the only speciated site with daily data coverage. Observed and simulated total and speciated PM<sub>2.5</sub> concentrations at the ATL site are decomposed into 9 or 10 IMFs (Fig. 3c). Because of the change in data frequency, 288 289 high frequency scales such as weekly cycles can be evaluated and the significance tested (Fig. S5) annual cycles with 290 the highest peak power is represented by IMF8 (IMF7 for SO<sub>4</sub> and NO<sub>3</sub>). Annual cycles of SO<sub>4</sub> and NO<sub>3</sub> appeared in 291 the earlier stage of decomposition in IMF7 because of their relatively weak half-year cycles, which largely led to the 292 mixed signal of half-year and annual cycles in IMF7 in total PM<sub>2.5</sub> as in Fig. 5b. This is more visible in the observed 293 IMF7 where the energy of the one-year period surpasses that of the half-year. Yet, clues can be seen from Fig. 5 that 294 the amplitude and the energy of annual cycles leaked into IMF7 is very limited compared to that remaining in IMF8, 295 indicating that it is still safe to conduct model evaluation on the seasonality using IMF8 with an underestimation in the amplitude of observation. On the other hand, inferences should be made with caution for IMF7 because of the 296 297 mixed modes. Scales up to 3 years are relatively well reproduced by the model.





## 300 **4.2 Long-term trend**

The EMD-decomposed long-term trend components for the observed and simulated total and speciated  $PM_{2.5}$ concentrations are presented in Fig. 6. To better visualize the non-linearity of the trend component, the rates of change (temporal derivative of a trend component, which is the change in the consecutive concentration divided by the sampling rate of 1 or 3 days and converted to the unit of  $\mu g/m^3/year$  by multiplying 365 day/year) are added with a separate y-axis on the right side in each panel (gray colored scale). It is evident that  $PM_{2.5}$  is changing at a varying rate, forming either a monotonic trend component or a trend component with one extremum, which cannot be fully represented by a single constant number using a traditional linear regression approach. Given that there are chemical

- 308 species (the remaining component, *Rem*) other than the ones studied in the total PM<sub>2.5</sub>, not all performance issues can
- 309 be fully explained by the five available species.

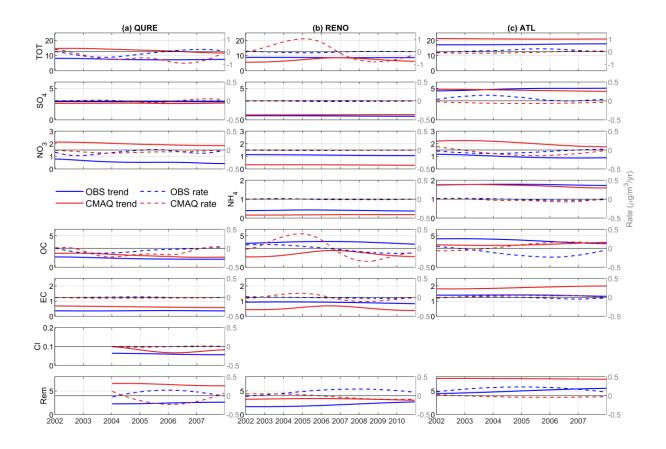


Fig. 6. Trend components of observed and simulated total and speciated PM<sub>2.5</sub> for (a) QURE, (b) RENO and (c) ATL in µg/m<sup>3</sup>. Dashed lines representing the rate of the change (temporal derivative of the trend component converted to µg/m<sup>3</sup>/year) are plotted against the right-side y axis, with a reference line of no change in black in the center. Note that the scales are not all the same.

315

316 At the QURE site, CMAQ captures the general decreasing trend in observed total PM2.5 which can mainly be traced 317 back to NO<sub>3</sub>, OC and the remaining components, while both observed and simulated trend components in SO<sub>4</sub> and EC 318 are relatively constant (Fig. 6a). The relative importance of each component in driving the trend of observed and 319 simulated total PM<sub>2.5</sub> reflected by its mean concentration share is summarized in Table 1 (time-dependent variations 320 of the concentration share is attached in Fig. S6 for reference). Moreover, the periods with highest decreasing rate in 321 observed total PM<sub>2.5</sub> during 2003-2004 with a decreasing rate of -0.44 µg/m<sup>3</sup>/year is also well replicated by the model. 322 Nevertheless, the slightly increasing  $PM_{2.5}$  level in the later years is simulated to be decreasing at a much higher rate, 323 which is partly due to the overestimated decreasing rate in OC and species other than the five studied ones. The trend 324 component of simulated Cl shows a cyclic-like feature because of proximity between the existence of a cycle of 4-5 325 years (by decomposing the simulation during the 6-year study period) and 4-year period limited by the available

- quality assured observations. The rate of change in the simulated trend component by decomposing the simulation
- during the 6-year study period would mimic that from the 4-year observation, both with a negligible negative value
- throughout 2004-2007. However, the mean magnitude of the trend component is almost twice as high (1.8 times
- 329 compared with observation) in the model with contribution from all species except for SO<sub>4</sub>. A quantitative summary
- of the comparison between the mean magnitudes of the observed and model trend components can be found in Table
- **331** 2.

**Table 1.** Concentration share (%) of different components in total  $PM_{2.5}$ . It is estimated by dividing the mean trend components of each species by that of total  $PM_{2.5}$  for both OBS and CMAQ, multiplied by 100. The concentration share of the remainder species *Rem* is estimated by subtracting all the available species share from 100 to compensate for the small discrepancies caused by the rounding up process and uncertainty in the mode decomposition. "-" indicates the data is not available (same applies for all other tables).

		$\mathbf{SO}_4$	$\mathbf{NO}_3$	$\mathbf{NH}_4$	OC	EC	Cl	Rem
QURE	OBS	38	7	-	19	5	1	30
	CMAQ	19	15	-	14	5	1	47
RENO	OBS	7	13	5	46	11	-	20
	CMAQ	11	4	2	30	7	-	45
ATL	OBS	28	6	10	24	8	_	24
	CMAQ	22	10	8	17	9	_	33

337

**Table 2.** The ratio of mean magnitude of the trend component  $r_{trend}$  (CMAQ/observation). Boldface values indicate a relatively good estimate of the magnitude (0.7 - 1.3).

	TOT	$\mathbf{SO}_4$	NO <sub>3</sub>	$\mathrm{NH}_4$	OC	EC	Cl
QURE	1.8	0.9	3.5	-	1.4	1.7	1.3
RENO	0.8	1.3	0.3	0.4	0.5	0.6	-
ATL	1.2	1.0	2.1	1.0	0.9	1.4	-

340

341

342 RENO is located close to the border with California and is affected by large wildfire breakouts in the western U.S. 343 (Gyawali et al., 2009) as can be seen in the spikes of the observed total PM<sub>2.5</sub> (Fig. 4a). Thus, OC makes up a much 344 larger portion of total PM<sub>2.5</sub> compared to other locations (Table 1). The model simulates large increasing rate up to 345 1.03  $\mu$ g/m<sup>3</sup>/year and decreasing rate up to -0.80  $\mu$ g/m<sup>3</sup>/year before and after the 2006-2007 winter season and fails to 346 reproduce the relatively stable condition seen in the observations with only  $-0.09 \ \mu g/m^3/year$  decreasing in 2004-2005 347 and 0.04 µg/m<sup>3</sup>/year increasing in 2008-2009 (Fig. 6b). Similar feature is found for combustion-related OC and EC 348 species. The observed slightly decreasing trends in SO<sub>4</sub> and NH<sub>4</sub> during 2005-2009 are not being captured in the model 349 simulations. The magnitude of the trend component is slightly underestimated with  $r_{trend}$  of 0.8 with contribution

from all species except for  $SO_4$  as well (Table 2).

- 351 During the period of 2002-2007, observations at ATL reveal a slightly increasing PM<sub>2.5</sub> trend that cannot be explained
- by the five available  $PM_{2.5}$  components trend (Fig. 6c), indicating a contribution of the remaining species such as the
- 353 non-carbonaceous portion of organic matter. Non-carbonaceous organic matter can account for more than half of total
- organic matter, which, in turn, can account for a large portion of the total PM<sub>2.5</sub> mass (Edgerton et al., 2005). In contrast,
- the model shows a slight decreasing trend with a peak decreasing rate in 2003 and misses the peak increasing rate of
- $0.23 \,\mu\text{g/m}^3/\text{year}$  in the winter season of 2005. Similarly, reversed trends are also simulated for SO<sub>4</sub>, OC and EC, while
- 357 the change rate in NO<sub>3</sub> is well captured. Unlike the previous sites, magnitude of trend components in total and
- speciated PM<sub>2.5</sub> are well simulated except for EC (1.4 times the observation) and NO<sub>3</sub> (2.1 times).
- 359 To sum up, the decreasing long-term trend at QURE is well simulated by the model. The occurrence of large wildfires
- 360 lasting for several months has significantly impacted the long-term trend component at RENO and the model failed
- 361 to capture those combustion-related species and total  $PM_{2.5}$  primarily due to limitations in the historical data used to
- 362 specify day-specific wildfire emissions (Xing et al., 2013). Slightly increasing levels of PM<sub>2.5</sub> and its species observed
- at ATL are simulated to be slightly decreasing, except for NO<sub>3</sub> which is well simulated. The magnitude of the long-
- term trend components of total PM<sub>2.5</sub> and SO<sub>4</sub> are well represented by CMAQ (Table 2). The model performs
- differently across the sites in terms of the magnitudes of the trend component in NO<sub>3</sub>, NH<sub>4</sub>, Cl, OC and EC. The large
- discrepancy in the magnitude of some long-term trend components is likely pointing to the systematic bias in the
- annual emission estimations as discussed in Xing et al., (2013), which mainly focused on long-term trend rather than
- the absolute level of the emissions. Species other than those in the available dataset also play a considerable role in
- 369 driving the agreements or disagreements between model simulations and observations of total PM<sub>2.5</sub>.

#### 370 **4.3 Seasonality**

371 The EMD-assisted seasonality evaluations utilize the decomposed IMFs with characteristic period of one year to 372 evaluate the amplitude and phase of the model simulation, both of which are time-dependent. As mentioned in Section 373 4.1, these IMFs are statistically significant from white noise with few exceptions (Fig. S5). We first demonstrate the 374 evaluation for total PM<sub>2.5</sub> at QURE (Fig. 7a). The top panel shows the annual cycle components and the bottom panel 375 shows its TDIC pyramid. The decreasing amplitude of the annual cycles throughout 2002-2007 is almost perfectly 376 represented with an overall ratio rannual being 1.0 (Table 3). Each pixel in the TDIC pyramid is the correlation (color-377 coded) calculated during a period of time I(t) with width of  $t_w$  days (y-axis) centered at a specific day (x-axis) as 378 introduced in Section 3.2. The annual cycle mean periods are identical between CMAQ and observations (350 days, 379 Fig. 2a IMF6), but there is a phase shift for all years with the entire TDIC pyramid being close to -1. By shifting the 380 CMAQ annual cycles backward 159 days (almost half year), the overall correlation of the annual component can reach 381 up to a peak of 0.9 (Table 4).

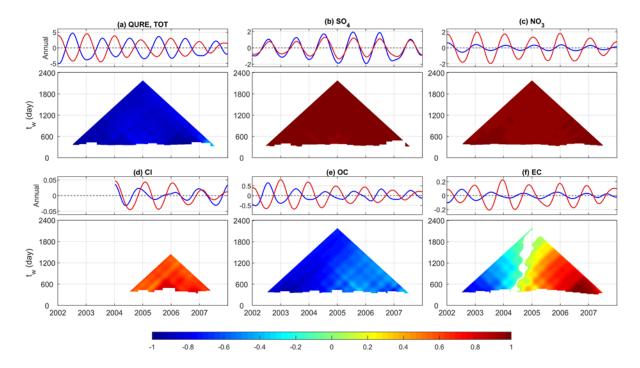
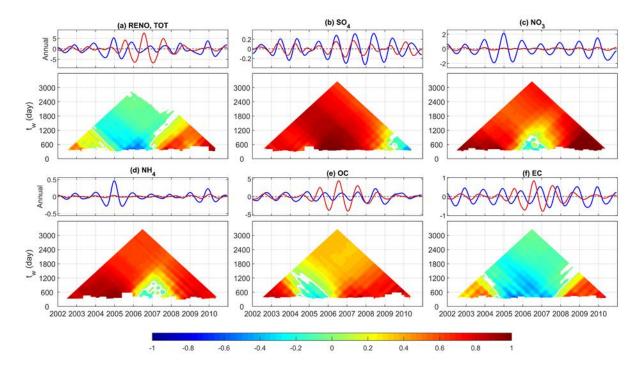


Fig. 7. Decomposed annual cycles (IMF6) from observed (blue) and simulated (red) concentrations ( $\mu$ g/m<sup>3</sup>) of (a) total PM<sub>2.5</sub>, (b) SO<sub>4</sub>, (c) NO<sub>3</sub>, (d) Cl, (e) OC and (f) EC and their corresponding TDIC at Quabbin Summit, MA. The window size  $t_w$  indicates the width of the window used to calculate a specific correlation centered at the day represented in x-axis.

387 What are the driving factors for the above phase shift in modeled total PM2.5 at Quabbin Summit, MA? The illustrations 388 in Fig. 7a for total  $PM_{2.5}$  alone cannot provide useful information that will allow the modeler to improve the model's 389 performance. This is accomplished by applying the EMD method to the PM<sub>2.5</sub> speciated components (Fig. 7b-f). Traces 390 of the semi-annual phase shift (-159 days) of annual cycles or large overestimation in the winter and underestimation in the summer is because of the largely overestimated amplitude of NO<sub>3</sub> (4.3 times that of observation) which peaks 391 392 in the winter and the almost semi-annual shifted OC (-147 days), as well as contributions from EC and Cl. NO<sub>3</sub> has a 393 mean amplitude reaching almost half of that of the total PM<sub>2.5</sub>. OC directly drives both the observed and simulated 394 annual components to be negatively correlated. EC follows the feature of OC in the first four years or so and the 395 feature of NO<sub>3</sub> in 2006 and 2007 and contributes to the half year shifted total PM<sub>2.5</sub>. The magnitude of winter-peaking 396 Cl cycles is overestimated with a phase shift of one month. However, the contribution of Cl is very limited because of 397 the tiny amplitude in both observed and simulated annual cycles. In addition, annual cycles in SO<sub>4</sub> are well reproduced 398 for the entire time span with an amplitude ratio of 0.7. A quantitative summary of the evaluation of the annual cycles 399 at this site can be found in Tables 3 and 4.



400

401 Fig. 8. Same as in Fig. 7 for Reno, NV, except that (d) represents NH<sub>4</sub> rather than Cl.

402 Both observed and simulated annual cycles at the RENO site are largely influenced by the extreme events lasting for 403 several months that are not properly simulated, indicating the need for more accurately specified wildfire emissions. 404 Overall, annual variations for total and speciated  $PM_{2.5}$  are largely underestimated except for the total  $PM_{2.5}$  and 405 combustion-driven EC and OC from 2005 to 2007 (Fig. 8). The modeled phase of SO<sub>4</sub>, NO<sub>3</sub>, NH<sub>4</sub> and OC agrees with 406 that of observation with the exception for a length of about two years in each that missed the phasing: 2009-2010 for 407 SO<sub>4</sub>, summer 2005-summer 2007 for NO<sub>3</sub>, 2006-2007 for NH<sub>4</sub> and 2004-2005 for OC. It is also notable that the TDIC 408 pyramid of EC mimics that of total PM<sub>2.5</sub>, implying the existence of errors in modeled EC in processes such as 409 emissions, transport, and deposition that affected the model performance for total PM2.5. In comparison, SO4 and OC 410 are relatively well simulated with a mean amplitude ratio of 0.5 and 1.5 and a phase shift of 36 and 33 days, 411 respectively.

412 Observed annual cycles of total PM<sub>2.5</sub> at the ATL site features a slightly increasing amplitude of annual variations 413 from 2002 to 2006 which then decreased to the original state in 2007 (Fig. 9a). Conversely, model-simulated annual 414 cycles became weaker throughout the period, with an overall  $r_{annual}$  of 0.5. As at the QURE site, the simulated annual 415 components at the ATL site also show a shift of several months (-132 days). Specifically, traces of these phase shifts 416 or large overestimation in the winter and underestimation in the summer can be seen from the more than doubled 417 amplitude of NO<sub>3</sub> which peaks in winter and underestimated SO<sub>4</sub> and NH<sub>4</sub> in the warm seasons as well as the -54 days 418 shifted EC. The anti-correlated remaining species other than those in the available dataset clearly played a role in 419 driving the discrepancies seen in the total PM<sub>2.5</sub> annual cycles (Fig. 10). Specifically, the anti-correlation likely points 420 to an inaccurate representation of the seasonal variation of the non-carbonaceous portion of organic matter due to an 421 incomplete representation of organic aerosols in the model version analyzed here; newer versions of the CMAQ model

- 422 include updated treatment of organic aerosols (e.g., additional SOA formation pathways, improvements in
- 423 representation of primary OM emissions) which is likely to correct the mentioned features (Appel et al., 2017; Murphy
- 424 et al., 2017; Xu et al., 2018). The underestimated annual variations in the remaining components closely resemble that
- 425 of the annual variation in total PM<sub>2.5</sub>. The phase of simulated SO<sub>4</sub>, NO<sub>3</sub>, NH<sub>4</sub>, and OC species is in good agreement
- 426 with those in observations and the amplitude of simulated annual cycles in SO<sub>4</sub>, OC and EC agree well with that in
- 427 the observations (Tables 3 and 4).
- 428 In sum, annual cycles of PM<sub>2.5</sub> are also time-dependent and the phase in the annual cycles for total PM<sub>2.5</sub>, OC and EC
- 429 reveals a general shift of up to half a year (Table 4); this indicates a potential problem in the allocation of emissions
- 430 during this study period and/or the treatment of organic aerosols in this version of the model. CMAQ generally
- 431 simulated the phase in SO<sub>4</sub>, NO<sub>3</sub>, Cl and NH<sub>4</sub> quite well but did not always capture the magnitude of their variations
- **432** (Table 3).

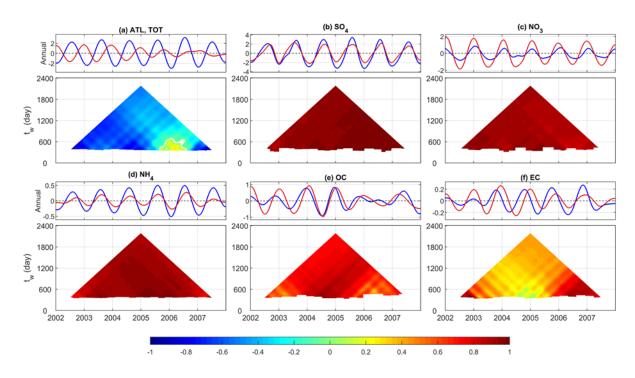
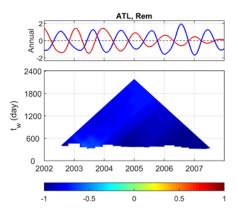


Fig. 9. Same as in Fig. 7 for Atlanta, GA, except that the annual component is resolved in IMF8 (IMF7 for SO<sub>4</sub>
and NO<sub>3</sub>) because of the difference in sampling rate and characteristic embedded in the time series at ATL and

436 (d) represents NH<sub>4</sub> rather than Cl.

437

433



# Fig. 10. Decomposed annual cycles in Atlanta, GA for the remaining components presented in total PM<sub>2.5</sub> other than the five species in Fig.9.

441 **Table 3.** The ratio of mean amplitude of the annual component  $r_{annual}$  (CMAQ/observation). Boldface values indicate

442 a magnitude with a ratio close to 1 (0.7 - 1.3).

	TOT	$SO_4$	NO <sub>3</sub>	$\mathrm{NH}_4$	OC	EC	Cl
QURE	1.0	0.7	4.3	-	1.6	3.1	1.6
RENO	1.2	0.5	0.1	0.2	1.5	0.9	-
ATL	0.5	0.7	2.4	0.4	1.2	1.0	-

443

**Table 4.** Phase shift (*n*) of CMAQ simulated annual cycle components in days. The background color indicates the maximum correlation ( $R_{max}$ ) that can be reached by shifting the CMAQ time series *n* days with respect to observations: white = [0.8, 1], light grey = [0.6, 0.8), grey = [0.4, 0.6), dark grey = (0.2, 0.4). The bold shows number of shifts less than a month while the italic shows shifts longer than three months.

	TOT	$SO_4$	NO <sub>3</sub>	NH <sub>4</sub>	OC	EC	Cl
QURE	-159	-6	3	-	-147	-105	-30
RENO	78	36	12	-21	33	96	-
ATL	-132	0	8	-17	-24	-54	-

448

#### 449 **4.4 Sub-seasonal and inter-annual variability**

450 In this section, model performance at multiple sub-seasonal and inter-annual scales with cycles less than 3 years,

451 presented in the total and speciated PM<sub>2.5</sub>, is evaluated following an approach similar to that for the annual cycles in

452 Section 4.3 (Fig. 11). First, IMFs from observations and model simulations are paired based on their characteristic

453 periods following the discussion in Section 4.1. Then, the magnitude of specific scales is evaluated using  $r_{IMFn}$ 

454 following Equation 6 of the  $r_{annual}$  for annual cycles. The phase shifts of the time series are assessed by the proportion

455 of shifted days relative to the mean characteristic scales of the corresponding observed and simulated IMFs  $\binom{n}{t_m}$ .

456 For example, a phase shift of 0.1 cycles in the 2-year cycles is approximately 73 days while it would be 18 days for457 the half-year cycles.

458 The performance of the simulated amplitude of the sub-seasonal and inter-annual cycles is relatively stable from a few 459 days to semi-annual scales and  $r_{IMFn}$  is close to 1 in most cases (Fig. 11a-c). CMAQ captures the features seen in the 460 observations at QURE, except for the large overestimation of NO<sub>3</sub> ( $r_{IMFn}$  ranges from 2.6 to 3.7 at the sub-seasonal 461 scale and reaches up to 13.8 for the 3-year cycles). Similar overestimation of NO<sub>3</sub> is also found at ATL ( $r_{IMFn}$  ranges 462 from 2.0 to 3.4, except for the 2-year cycles). In contrast, NO<sub>3</sub> at RENO is strongly underestimated with  $r_{IMFn}$  ranging 463 from 0.1 to 0.3 and reaching its minimum at the 2-year cycles. Likewise, all time scales of NH<sub>4</sub> at RENO are also 464 being underestimated with  $r_{IMFn}$  decreasing from 0.4 to only 0.1 at the 3-year cycles. The coexistence of 465 underestimation of NO<sub>3</sub> and NH<sub>4</sub> variability, as well as their trend component, likely points to the insufficient grid 466 resolution in representing ammonium nitrate episodes associated with stagnant meteorology in the mountainous 467 regions as illustrated by Kelly et al. (2019). To sum up, model has simulated the magnitude of features across all scales 468 in most of the studied cases. However, fluctuations in NO<sub>3</sub> are constantly being largely over- or under-estimated and

469 improvements to the model are required to better replicate its variability (Fig. 11a-c).

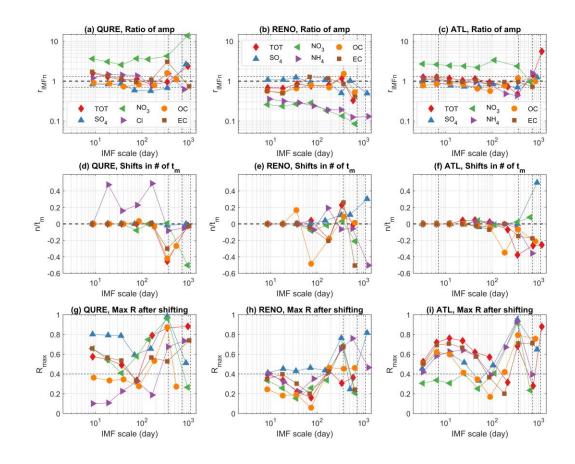
470 A high  $R_{max}$  of corresponding IMFs can only be achieved when the characteristic scales of those from observations 471 and model simulations are close, there is minimal mode mixing, and negligible irregular change of amplitude exists 472 during the study period. Thus,  $R_{max}$  tends to be small for all oscillations at RENO because of the irregular impact 473 from events such as wildfires. Thus, the interpretation of phase shift is focused on the components and time scales 474 having correlations above 0.4 only.

475 Results show that the sub-seasonal cycles at QURE all have a negligible phase shift of less than 0.1 cycles (Fig. 11d). 476 The semi-annual cycles at RENO have around 0.2 cycle phase shifts in total  $PM_{2.5}$  (-0.2),  $NH_4$ (0.2), OC (-0.2), and 477 EC (-0.2) while negligible phase shifts of less than 0.1 cycles are simulated in SO<sub>4</sub> ranging from 9 days to semi-annual 478 in scale. As at QURE, multiple sub-seasonal cycles at ATL all have a negligible phase shift of less than 0.1 cycles, 479 with the exception of semi-annual OC which has a phase shift of nearly -0.4 cycles with a marginal correlation of 480 around 0.4. Unlike the relatively stable  $R_{max}$  throughout the time scales within each of the species for QURE and 481 RENO, the  $R_{max}$  at ATL tends to be much higher (roughly 0.6-0.8) in the scales of 6 to 25 days, except for NO<sub>3</sub>, 482 indicating the model's success in simulating those weather-induced air quality fluctuations at this site as reflected by 483 their negligible phase shifts.

484 However, the physical meaning of each sub-seasonal IMF is not yet fully understood and requires further study. For 485 example, synoptic scale IMFs (IMFs with scale less than/around a month) usually have large variance and are not 486 statistically significantly different from white noise except for observed SO<sub>4</sub> and NH<sub>4</sub> (Fig. S5). Yet, observed and 487 simulated total and some speciated PM<sub>2.5</sub> at QURE and ATL (except IMF1) can achieve moderate to high  $R_{max}$  at 488 these time scales (Fig. 11 g-i), indicating a potential physical explanation of those time scales using meteorological 489 variables. IMFs with scales longer than a month but less than half year possess much less variance and are usually not

490 statistically significantly different from noise. Exceptions are also found at the Atlanta site where observed IMFs are

- 491 mostly significantly different from noise. Whereas semi-annual cycles are mostly statistically significant (note that
  492 semi-annual SO<sub>4</sub> and NO<sub>3</sub> at ATL are too weak to be decomposed into a separate IMF). In a previous study, He et al.
  493 (2014) found semi-annual oscillations in the corrected AErosol RObotic NETwork (AERONET) Aerosol Optical
- 494 Depth (AOD) and PM<sub>10</sub> mass concentrations are primarily caused by the change of wind directions in Hong Kong.



496

497 Fig. 11. Model performance at all temporal scales for sites QURE, RENO and ATL. (a-c) ratio of mean 498 amplitude of corresponding IMFs with similar characteristic mean periods (ideal ratio=1.0); (d-f) the phase 499 shift n in the number of mean periods (average mean period of corresponding IMFs decomposed from observation and model simulation); (g-i) maximum correlation  $R_{max}$  can be achieved by shifting the modeled 500 501 time series. The average mean period of corresponding IMFs decomposed from observations and CMAQ of 502 total and speciated PM2.5 are represented on the x-axis; all metrics on the y-axis are unitless. Horizontal 503 reference lines are drawn at 0.7 and 1.3 in (a-c). Weekly, annual and inter-annual (2- to 3-year) scales are 504 marked with vertical dashed lines.

505 The evaluation and interpretation of inter-annual cycles are constrained by the limited available speciated observations

- 506 for a period of 6 to 9 years (4 years for Cl at QURE). Thus, only 2- to 3-year cycles are presented (Fig. 11) and
- 507 evaluated. Among the 2- to 3-year inter-annual cycles at QURE, there is minimal phase shift for total PM<sub>2.5</sub>, SO<sub>4</sub>, Cl,
- and EC with moderate to high  $R_{max}$ . At RENO, the model presents negligible shifts in 2-year cycles of OC and NH<sub>4</sub>

509 while phase shifts of 0.3 and -0.5 cycles are simulated in the 3-year cycles for SO<sub>4</sub> and NH<sub>4</sub>. At ATL, the phase shift

510 of -0.2 to -0.4 cycles are simulated for PM<sub>2.5</sub>, NH<sub>4</sub>, OC and EC with periods of 2- to 3-year cycles; while 2- to 3-year

511 SO<sub>4</sub> cycles have a half-year cycle shift.

#### 512 **5 Conclusions**

513 The main advantage for using EMD to evaluate PM<sub>2.5</sub> and its speciated components is that it decomposes nonlinear 514 and nonstationary signals into multiple modes and a residual trend component. It does not require any preselection of 515 the temporal scales and assumptions of linearity and stationarity for the data, thereby providing insights into time 516 series of PM<sub>2.5</sub> concentrations and its components. Using improved CEEMDAN, we are able to assess how well 517 regional-scale air quality models like CMAQ can simulate the intrinsic time-dependent long-term trend and cyclic 518 variations in daily average PM<sub>2.5</sub> and its species. This type of coordinated decomposition and evaluation of total and 519 speciated  $PM_{2.5}$  provides a unique opportunity for modelers to assess influences of each  $PM_{2.5}$  species to the total 520 PM<sub>2.5</sub> concentration in terms of time shifts for various temporal cycles and the magnitude of each component including 521 the trend.

- 522 A demonstration of how improved CEEMDAN could be applied to PM<sub>2.5</sub> time series at three sites over CONUS that 523 provide speciated PM2.5 data reveals the presence of the annual cycles in  $PM_{2.5}$  concentrations and time-dependent 524 features in all decomposed components. At these three sites, the model generally is more capable of simulating the 525 change rate in the trend component than the absolute magnitude of the long-term trend component. However, the 526 magnitude of SO<sub>4</sub> trend components is well represented across all three sites. Also, the model reproduced the amplitude 527 of the annual cycles for total PM<sub>2.5</sub>, SO<sub>4</sub> and OC. The phase difference in the annual cycles for total PM<sub>2.5</sub>, OC and 528 EC reveal a shift of up to half-year, indicating the need for proper allocation of emissions and an updated treatment 529 of organic aerosols compared to the earlier model version used in this set of model simulations. The consistent large 530 under/over-prediction of NO<sub>3</sub> variability at all temporal scales and magnitude in the trend component, as well as the 531 abnormally low correlations of synoptic scale NO<sub>3</sub> at ATL, calls for better representation of nitrate partitioning and 532 chemistry. Wildfires have the potential to elevate  $PM_{2.5}$  for months and can alter its variability at scales from few days 533 to the entire year. Thus, more accurate fire emission data should be incorporated to improve model simulation, 534 especially in those fire-prone regions.
- 535 Data availability. Paired observations and CMAQ model data used in the analysis will be made available at
  536 https://edg.epa.gov/metadata/catalog/main/home.page. Raw CMAQ model outputs are available on request from the
  537 U.S EPA authors.
- Author contribution. "HL and MA designed the methodology; RM, CH and SR contributed in the assessment of the
  outcomes and were consulted on necessary revisions. Model simulations were performed by the US EPA. HL prepared
  the manuscript with contributions from all co-authors."
- 541 Acknowledgements

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