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Characteristics and sources of aerosol aminiums over the eastern coast of China: Insights from the integrated observations in a coastal city, adjacent island and the marginal seas

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Abstract. An integrated observation on aerosol aminiums was conducted in a coastal city (Shanghai) of eastern China, a nearby island (Huaniao Island) and over the Yellow Sea and East China Sea (YECS). Triethylaminium ($TEAH^+$) was the most abundant aminium observed in Shanghai but not detected over the island and the open seas, suggesting its predominantly terrestrial origin. By contrast, relatively high concentrations of dimethylaminium ($DMAH^+$) and trimethylaminium+diethylaminium ($TMDEAH^+$) were measured over the ocean sites. Environmental factors, including boundary layer height (BLH), temperature, atmospheric oxidizing capacity and relative humidity, were found to be related to aminium concentrations. All the detected aminiums demonstrated the highest levels in winter in Shanghai, consistent with the lowest BLH, temperature and oxidizing capacity in this season. Aminiums mainly existed in fine particles and showed a bimodal distribution with two peaks at 0.18–0.32 μm and 0.56–1.0 μm, indicating that condensation and cloud processing were primary formation pathways for aminiums. Nonetheless, a unimodal distribution for aerosol aminiums was usually measured over the YECS or influenced mainly by the marine air-mass over the Huaniao Island, which was probably related to sea-spray aerosols that either contained primary aminiums or provided surface for heterogeneous reactions to form secondary aminiums. Terrestrial anthropogenic sources and marine biogenic sources were both important contributors for $DMAH^+$ and $TMDEAH^+$, and the latter exhibited a significantly higher $TMDEAH^+$ to $DMAH^+$ ratio. By using the mass ratio of methanesulfonate (MSA) to non-sea-salt SO_4^{2-} as an indicator of marine biogenic source, we estimated that marine biogenic source contributed to 57–83% and 29–38% of aerosol aminiums over Huaniao Island in the summer of 2017 and autumn of 2016, respectively.

1 Introduction

Low molecular weight amines are commonly found in the atmosphere in both gaseous and particulate phases (Ge et al., 2011b, a). Base on present theoretical calculations (Kurten et al., 2008; Loukonen et al., 2010; Paasonen et al., 2012; Olenius et al., 2017), laboratory simulations (Wang et al., 2010a; Wang et al., 2010b; Kurten et al., 2014; Erupe et al., 2011; Almeida et al., 2013; Yu et al., 2012) and field observations (Smith et al., 2010; Kürten et al., 2016; Tao et al., 2016), amines in the atmosphere have been proved to play an important role in new particle formation and subsequent particle growth, and thus affect both the number concentrations of aerosols and cloud condensation nuclei which are closely relevant to regional climate (Tang et al., 2014; Yao et al., 2018). For example, dimethylamine (DMA) was found to be a key species involved in new particle formation events in the urban area of Shanghai, and the nucleation mechanism was likely to be H_2SO_4-DMA-H_2O ternary nucleation (Yao et al., 2018). Gaseous amines in the atmosphere can react with oxidants such as ·OH and O_3 to form secondary organic aerosols (SOA) (Murphy et al., 2007) or other gases (Nielsen et al., 2012). The heterogeneous reaction, such as replacing the

NH_4^+ in particles, is another important pathway for amines to form SOA in the atmosphere (Pankow, 2015; Kupiainen et al., 2012; Liu et al., 2012; Chan and Chan, 2013). In aerosols, amines are mainly in the form of protonated cations, namely aminiums (Ge et al., 2011a).

Amines originate from a wide range of sources, including anthropogenic sources such as animal husbandry and industrial emissions, as well as natural sources such as marine sources, vegetation emissions, and soil processing, etc. (Ge et al., 2011b; Hemmilä et al., 2018). Zheng et al. (2015) measured amines in a suburban site of Nanjing in China, and concluded that amines and NH_3 in the region were mainly from industrial emissions in adjacent areas. Shen et al. (2017) demonstrated that coal combustion could emit abundant methylaminium ($MMAH^+$), ethylaminium ($MEAH^+$) and diethylaminium ($DEAH^+$) through combustion experiments, and the corresponding emission factors were 18.0±16.4, 30.1±25.6 and 14.6±10.1 mg kg^{-1}, respectively. In marine boundary layer, marine source is an important contributor for amines and it is closely related to marine surface biological activities. In the North Atlantic, the concentrations of dimethylaminium ($DMAH^+$) and $DEAH^+$ were significantly higher during the periods with high biological activity and clean air-mass conditions than those with low biological activity or polluted air masses advecting to the sampling site, and the contributions of these two aminiums to SOA and water soluble organic nitrogen (WSON) reached 11% and 35%, respectively (Facchini et al., 2008). The observation in Cape Verde also showed that the concentrations of amines were higher during the occurrence of algal blooms (Müller et al., 2009). Previous studies on aminiums over the marginal seas of China indicated that $DMAH^+$ and trimethylaminium ($TMAH^+$) were overwhelmingly from marine sources (Hu et al., 2015; Yu et al., 2016; Xie et al., 2018). In May 2012, the concentrations of $DMAH^+$ and $TMAH^+$ over the Yellow Sea (YS) and Bohai Sea even reached 4.4±3.7 and 7.2±7.1 nmol m^{-3}, which was 1–3 orders of magnitude higher than those reported in other oceanic regions (Hu et al., 2015). These extremely high concentrations were thought to be associated with high biological activities.

Given the potentially important roles of amines in the atmosphere and the complexity of their sources, it is important to conduct a systematic analysis on their concentrations, affecting factors, formation pathways and source contributions. The eastern China is a densely populated region with strong human activities and large emissions of atmospheric pollutants. Under the influence of the summer monsoon, marine source components can be vital to the atmospheric composition of the coastal area. Although the lifetime of gaseous amines in the atmosphere is only a few hours, it can be prolonged after amines partition into the particulate phase, and thus, they may be transported over a long range (Nielsen et al., 2012). Many studies have been done on the atmospheric amines over eastern China and adjacent seas (Huang et al., 2012; Hu et al., 2015; Zheng et al., 2015; Huang et al., 2016; Tao et al., 2016; Yu et al., 2016; Shen et al., 2017; Xie et al., 2018; Yao et al., 2018; Yao et al., 2016). Nonetheless, the long-term observation of aminiums over the coastal sea and quantitative estimate of the contribution of marine biogenic source to aerosol aminiums are still lacking.

In this study, the aminiums over a coastal megacity (Shanghai), a nearby island (Huaniao Island) and marginal seas (the Yellow Sea and East China Sea, YECS) were measured. The relationships between aminium concentrations and environmental factors were systematically analyzed. The size distributions of aminiums were investigated with the speculation of primary formation pathways. Besides, the dominant sources determining the concentrations and ratios between aminium species were elucidated, and the contributions of terrestrial anthropogenic and marine biogenic sources to aminiums were quantitatively estimated. Our results will be a great help for understanding the chemical properties, reaction pathways and sources of aerosol aminiums over the coastal area and the ocean.

2 Sampling and Analysis

2.1 Aerosol sampling

The sampling site in Shanghai was located on top of the No.4 teaching building of Fudan University (31.30° N, 121.50° E) (Fig. 1). This site is affected by the school, residential, commercial and traffic activities and can be a representative of coastal

cities. Particulate matters with an aerodynamic diameter less than 2.5 μm ($PM_{2.5}$) were simultaneously collected by two medium-flow samplers (100 L min^{-1}, HY-120B, Hengyuan) using a 90 mm pre-combusted quartz filter (Whatman) and a cellulose filter (Grade 41, Whatman), respectively. A total of 131 samples were collected within four seasons with the sampling duration ~24 hours (Table 1).

Aerosols were also collected at Huaniao Island (HNI, 30.86° N, 121.67° E) which was about 80 km away from Shanghai in the East China Sea (ECS) (Fig. 1). The locally anthropogenic emissions were negligible, but the site was affected by the terrestrial transport and the ship emission from nearby container ports (Wang et al., 2016; Wang et al., 2018). Fourteen $PM_{2.5}$ samples were collected in the summer of 2016 and size-segregated samples were obtained using a 10-stage Micro-Orifice Uniform Deposit Impactor (30 L min^{-1}, MOUDI, MSP Model 110-NR) and 47 mm PTFE filters (Zeflour, PALL) between 2016 fall and 2017 late summer (Table 1). The 50% cutoff diameters for 10 stages were 18, 10, 5.6, 3.2, 1.8, 1.0, 0.56, 0.32, 0.18, 0.10 and 0.056 μm, and the sampling durations were 24-48 hours.

The size-segregated samples were also collected over the YECS onboard research vessel (R/V) *Dong Fang Hong II* in the spring of 2017. The cruise started from Qingdao on March 27 and returned on April 15 (Fig. 1), and a total of 9 sets of samples were obtained.

2.2 Chemical analysis

One fourth of $PM_{2.5}$ sample and half of MOUDI sample filters were cut and placed into a polypropylene jar (Nelgene) with 20 mL of ultrapure water (18.25 MΩ cm^{-1}) for 40 min ultrasonic extraction. The extract was filtered through a 0.45 μm PTFE filter (Jinteng) and stored at 4 °C for ion measurement. Ion Chromatograph (DIONEX ICS-3000, Thermo-Fisher) assembled with AG11-HC and AS11-HC was used to determine anions including Cl^-, NO^{3-}, SO_4^{2-}, $HCOO^-$, methanesulfonate (MSA), malonate, succinate, glutarate, maleate and $C_2O_4^{2-}$. The columns CG17 and CS17 were used to measure inorganic cations including Na^+, NH_4^+, K^+, Mg^{2+} and Ca^{2+} and aminiums. The detailed procedures for meusuring $DMAH^+$, $TMAH^+$+$DEAH^+$, propylaminium ($MPAH^+$), triethylaminium ($TEAH^+$), ethanolaminium ($MEOAH^+$) and triethanolaminium ($TEOAH^+$) refer to Zhou et al. (2018). It should be noted that $TMAH^+$ and $DEAH^+$ could not be completely separated using the IC system (VandenBoer et al., 2012; VandenBoer et al., 2011; Zhou et al., 2018; Huang et al., 2014). Nonetheless, the sum of $TMAH^+$ and $DEAH^+$ concentrations (referred to $TMDEAH^+$) might be quantified using the calibration curve of $TMAH^+$ with errors less than 3% (Zhou et al., 2018).

One fourth of $PM_{2.5}$ cellulose sample filter was cut and digested with 7 mL of HNO_3 and 1 mL of HF (both acids were purified from GR using a sub-boiling system) at 185 °C for 30 min in a microwave digestion system (MARS5 Xpress, CEM). An Inductively Coupled Plasma Optical Emission Spectroscopy (ICP-OES, SPECTRO) was used for determining elements Al, Ca, Fe, Na, P, S, Cu, K, Mg, Mn, Zn, As, Ba, Cd, Ce, Co, Cr, Mo, Ni, Pb, Ti, and V. The detailed procedures refer to Wang et al. (2016).

2.3 Auxiliary data

The 3-hour resolution meteorological data of Baoshan station in Shanghai (WMO index: 58362) were obtained from the National Climatic Data Center (NCDC, https://www.ncdc.noaa.gov/isd). The 10-second resolution meteorological data were recorded by a shipborne meteorological station during the cruise. The planetary boundary layer height (BLH) and 6-hour accumulated precipitation (TPP6) for the cruise were extracted from NCEP's Global Data Assimilation System Data (GDAS). The daily concentrations of gaseous pollutants (SO_2, CO, NO_2 and O_3) in Shanghai were obtained from the Shanghai Environmental Monitoring Center (http://www.semc.gov.cn/aqi/home/DayData.aspx).

Three-day air mass backward trajectories were calculated using a Hybrid Single-Particle Lagrangian Integrated Trajectory (HYSPLIT) model (http://ready.arl.noaa.gov/HYSPLIT.php) with the starting height at 100 meters.

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3 Results and discussion

3.1 Seasonal and spatial variations of aminium concentrations

Three aminiums, $DMAH^+$, $TMDEAH^+$ and $TEAH^+$, were commonly detected in the aerosol samples collected from Shanghai. The most abundant aminiums were $DMAH^+$ and $TEAH^+$ with their annual means of 15.6 and 16.0 ng m^{-3}, respectively. By comparison, the average $TMDEAH^+$ concentration (4.4 ng m^{-3}) was significantly lower. All three aminiums showed the highest concentrations in winter and the lowest levels in spring ($DMAH^+$) and summer ($TMDEAH^+$ and $TEAH^+$), which generally agreed with the seasonal trends of $PM_{2.5}$ and NH_4^+ concentrations in Shanghai (Figure 2). Specifically, the average $TEAH^+$ reached 35.2 ng m^{-3} in winter in Shanghai, about 40 times as much as that in summer. By contrast, $TEAH^+$ was mostly below the detection limit in the aerosols collected over Huaniao Island and the YECS, suggesting its dominant land sources and negligible marine contribution. Differently, the average $DMAH^+$ and $TMDEAH^+$ concentrations (14.0 and 13.2 ng m^{-3}) over Huaniao Island were close to and significantly higher than those of Shanghai, respectively. Similarly high concentrations of $DMAH^+$ and $TMDEAH^+$ (11.9 and 14.6 ng m^{-3}) were also observed over the YECS (Fig. 2 and Table 2), suggesting that the two aminiums might have notable marine sources. Accordingly, both species reached the highest levels during the summer campaigns in 2017 at Huaniao Island, consistent with the highest primary productivity in the coastal ECS and prevailing winds from the ocean in summer. As a major component of fine particles over eastern China with similar chemical properties to aminiums, NH_4^+ was mainly from terrestrial sources and its concentrations over Huaniao Island were much lower than those over Shanghai (Fig. 2).

Our measurement of $DMAH^+$ in Shanghai was comparable to those previously reported from the urban sites (Table 2), but generally higher than those measured in the forest areas of Toronto (VandenBoer et al., 2012), Hyytiälä (Hemmilä et al., 2018) and Guangdong (Liu et al., 2018a). This implies that anthropogenic activities may be crucial sources of $DMAH^+$ in the urban atmosphere. The $TMDEAH^+$ concentrations in our study were much lower than those reported by Tao et al. (2016) in Shanghai. Their sampling location was close to the residential areas and could be influenced by the local sources such as human excreta emission (Zhou et al., 2018). The aerosol $TEAH^+$ concentrations in China were firstly reported in our study and could not be compared to previous work. Except for the three aminiums, $MMAH^+$ and $MEAH^+$ (Liu et al., 2018a; Ho et al., 2015; Shen et al., 2017) were other abundant aminiums detected in the urban site.

Aerosols were sampled using a MOUDI over Huaniao Island and the YECS. Aminiums in $PM_{1.8}$ of the MOUDI samples were compared to those of $PM_{2.5}$, since MOUDI does not have the 50% cutoff diameter of 2.5 μm and aminiums in $PM_{1.8}$ accounted for over 60% concentrations of the whole size range of aerosols. Our measurements of aminiums over Huaniao Island and the YECS were comparable to those previously observed over the eastern China seas (Hu et al., 2015; Yu et al., 2016; Xie et al., 2018), but they were apparently higher than many other oceanic regions such as Arabian Sea (Gibb et al., 1999) and Cape Verde (Müller et al., 2009). The high aminiums over the YECS were probably associated with the severe air pollution in eastern China as well as the high ocean productivity in marginal seas.

3.2 Environmental factors affecting aminium concentrations

3.2.1 Boundary layer height (BLH)

The concentrations of $PM_{2.5}$, NH_4^+ and three aminiums sampled in Shanghai in 2013 dropped significantly when the BLH increased from 200 m to 500 m and then slowly decreased with the further increase of BLH (Fig. 3a and Fig. S1), due to the improvement of diffusion condition. Specifically, the concentrations of $DMAH^+$, $TMDEAH^+$ and $TEAH^+$ (58.4, 13.9 and 80.5 ng m^{-3}) in Shanghai reached the maximum along with $PM_{2.5}$ (447 μg m^{-3}) during the severe haze event between 30 Nov. and 8 Dec. 2013, when the average BLH and wind speed were 298 m and 1.35 m s^{-1}, respectively (Fig. S2). By comparison, the average concentrations of $DMAH^+$, $TMDEAH^+$ and $TEAH^+$ (8.9, 4.0 and 10.1 ng m^{-3}) were much lower prior to the haze event (on 26-29 Nov 2018) associated with the higher BLH (636.4 m) and wind speed (2.73 m s^{-1}). Thus, the generally poor diffusion

condition in winter (Liu et al., 2013) could cause a substantial increase of aminiums in aerosols and lead to the seasonal variation of aminiums in Shanghai.

3.2.2 Temperature

To eliminate the synchronous change of aminums and NH_4^+ with $PM_{2.5}$, the mass ratios of aminiums to $PM_{2.5}$ (aminiums/$PM_{2.5}$) and NH_4^+ to $PM_{2.5}$ ($NH_4^+/PM_{2.5}$) were applied for analysis. These ratios were found to be negatively correlated with air temperature in Shanghai (Fig. 3b). Similar to NH_4^+, aminiums combined with NO_3^-, Cl^- and organic acids are semi-volatile and can dissociate in the atmosphere (Tao and Murphy, 2018). So the negative correlations may be explained by the movement of gas-particle partitioning equilibrium to the gas phase at higher temperatures (Ge et al., 2011a). This is consistent with the previous observation that the proportion of particles containing aminiums in the urban area of Shanghai was much higher in winter (23.4%) than that in summer (4.4%) (Huang et al., 2012). The seasonal variation of temperature may also lead to the change of concentrations of aerosol aminiums.

3.2.3 Oxidizing capacity

As gaseous amines can be oxidized by oxidants such as ·OH, O_3 and NO_3· in the atmosphere before partitioning into the particulate phase (Ge et al., 2011b; Nielsen et al., 2012; Yu and Luo, 2014), aminium concentrations in aerosols may decrease with the enhanced atmospheric oxidizing capacity. Ozone concentration can represent oxidizing capacity of the lower atmosphere (Thompson, 1992). Here the relationship between aminium/NH_4^+ ratios and O_3 was examined, because the formation of particulate aminiums and NH_4^+ were both temperature-dependent and using their ratios could avoid the temperature effect to some extent. Besides, the residence time of NH_3 in the atmosphere due to the oxidation reaction is about 72.3 days (Ge et al., 2011b), and therefore NH_4^+ concentrations in aerosols should not be affected by O_3. A negative correlation was found between the $TEAH^+/NH_4^+$ and O_3 concentrations in Shanghai (Fig. 3c). Differently, the $DMAH^+/NH_4^+$ and $TMDEAH^+/NH_4^+$ reached the highest values at the mid-level O_3 and decreased with both low and high concentrations of O_3. This verifies that high oxidizing capacity may reduce the formation of particulate aminiums by oxidizing gaseous amines. This also implies that $DMAH^+$ and $TMDEAH^+$ may have the sources different from $TEAH^+$ but similar to O_3 precursors such as biogenic VOCs. Among the three amines, the rate constants of TEA reacting with ·OH and O_3 were larger than those of other two amines (Nielsen et al., 2012), and thereby $TEAH^+$ showed the most significant correlation with O_3. In general, atmospheric oxidizing capacity was the strongest in summer (Logan, 1985; Liu et al., 2010), which could be another reason for seasonal variation of aerosol aminiums in Shanghai.

In the spring of 2017 over the YECS, the concentrations of $DMAH^+$ and $TMDEAH^+$ were found to be the lowest between 29 Mar and 4 Apr when it was sunny and Chl-a concentrations were relatively low. The relatively low biogenic emission may partly account for the low-level aminiums. Nonetheless, the $HCOO^-$ in aerosols, a product of photochemical reactions under high oxidizing capacity (Souza, 1999; Tsai et al., 2013), reached the highest level between 31 Mar. and 4 Apr. (42.1–55.5 ng m^{-3}). Its concentrations were inversely correlated with aminiums when eliminating the lowest values of $HCOO^-$ (Fig. 4). This further suggests that high oxidizing capacity may be one of causes for lowered aminiums in marine aerosols.

3.2.4 Relative humidity and fog processing

In the spring of 2017 over the YECS, although the sample of 4–5 Apr. was influenced by high Chl-a concentrations and low BLH, the concentrations of $DMAH^+$ and $TMDEAH^+$ (13.3 and 17.4 ng m^{-3}) were about half of those on 7–9 Apr. (Fig. 5). This was probably due to the intense fog event occurred on 7–9 Apr. with relative humidity>90%, which could enhance the gas-to-particle partitioning of amines. The enhancement of TMA gas to particles by cloud and fog processing has been observed in both field and laboratory simulations (Rehbein et al., 2011). It was also found that the number fraction of TMA-containing particles dramatically increased from ~7% in clear days to ~35% in foggy days and number-based size distribution of TMA-

containing particles shifted towards larger mode, peaking at the droplet mode (0.5–1.2 μm) in Guangzhou (Zhang et al., 2012). The investigation over the Yellow and Bohai seas in the summer of 2015 found significantly positive correlations between the concentrations of $DMAH^+$ and $TMAH^+$ and relative humidity (Yu et al., 2016). Therefore, high relative humidity and fog event may lead to an increase of aminiums in marine aerosols.

3.3 Size distributions and formation pathways of aerosol aminiums

The aminiums were mainly distributed in fine aerosols with diameter less than 1.8 μm, and the mass percentages of $DMAH^+$ and $TMDEAH^+$ in the coarse mode were around 36% in the autumn of 2016 at Huaniao Island and less than 15% in all other campaigns at Huaniao Island and over the YECS (Fig. 6a-d). The aminiums mostly demonstrated a bimodal distribution in the autumn and early summer campaigns at Huaniao Island with peaks at 0.18–0.32 μm (condensation mode) and 0.56–1.0 μm (droplet mode). This is similar to the size distributions of $DMAH^+$ and $TMDEAH^+$ observed in Shanghai (Tao et al., 2016) and to NH_4^+ and non-sea-salt ($nss\text{-}SO_4^{2-}$) in all campaigns over Huaniao Island and the YECS (Fig. S3-4). The size distribution suggests that the gas-to-particle condensation (condensation mode) and cloud processing (droplet mode) seem to be primary mechanisms for the formation of aminiums and other secondary species NH_4^+ and $nss\text{-}SO_4^{2-}$.

In order to compare the contributions between condensation and cloud processing to the formation of specific species, the ratio of its concentrations in droplet mode (0.56–1.0 μm) to condensation mode (0.18–0.32 μm) was calculated (denoted as α). It could be seen that the α values of NH_4^+ and $nss\text{-}SO_4^{2-}$ were significantly greater than 1, especially in the case of high concentrations, indicating that the cloud processing probably determined the concentrations of these species (Fig. 7). Differently, aminiums had α values around 1, suggesting that condensation and cloud processing might be equally important to the formation of aminiums.

In late summer at Huaniao Island and the spring cruise over the YECS when air masses were mainly from oceanic regions (see Sect. 3.4.3), the aminiums generally exhibited a unimodal distribution with one wide peak at 0.18–1.0 μm due to the increased concentrations at 0.32–0.56 μm (Fig. 6e-h). The concentrations of NH_4^+ and $nss\text{-}SO_4^{2-}$ also showed a significant elevation in the size range of 0.32–0.56 μm during these periods. The deviation of MOUDI cutoff diameters during the sampling could be ruled out because the concentrations of particulate matter always presented a trimodal distribution with peaks at 0.18–0.32 μm, 0.56–1.8 μm and 3.2–10 μm. The unimodal distributions of aminiums with the peak at 0.18–1.0 μm have been widely reported over the eastern China seas (Hu et al., 2015; Yu et al., 2016; Xie et al., 2018). This suggests that the formation mechanisms of aerosol aminiums over the ocean may be different from that in the urban area. It was indicated that the high concentration and unique size distribution of $TMAH^+$ observed over the oligotrophic western North Pacific were mainly attributed to the primary $TMAH^+$ in sea-spray aerosols (Hu et al., 2018). So we speculate that the elevated concentrations of aminiums at 0.32–0.56 μm over the eastern China seas may be also associated with the increased concentration of sea-spray aerosols which contain substantial primary aminiums or provide more surface for heterogeneous reactions to form secondary aminiums (Yu et al., 2016).

3.4 Sources of aerosol aminiums

3.4.1 Anthropogenic sources on land

Correlation analysis was carried out between aminiums, other $PM_{2.5}$ components and gaseous pollutants measured in Shanghai (Fig. 8). It can be seen that the secondary inorganic components SO_4^{2-}, NO_3^- and NH_4^+ (SNA), $PM_{2.5}$ and $DMAH^+$ were significantly correlated with each other with the correlation coefficients above 0.6. This suggests that anthropogenic sources may have a great contribution to the atmospheric DMA in Shanghai. The correlations between $TEAH^+$ and SNA were relatively weak, but $TEAH^+$ was found to be significantly correlated with the components mainly from industrial sources (represented by the high concentrations of K, Mn, Cd, Pb, Zn, and Cl^-) (Tian et al., 2015; Liu et al., 2018b), indicating that the industrial

emission could be an important source of TEA. Compared to the $DMAH^+$ and $TEAH^+$, $TMAH^+$ showed much weaker correlations with the anthropogenically derived components. Weak correlations were also found between all the aminiums and V, Ni, Al, Mg, Ca and Fe, suggesting that ship emission (traced by V and Ni) and soil dust (represented by Al, Ca and Fe) were not main sources of aminiums in $PM_{2.5}$ over Shanghai.

3.4.2 Marine biogenic source

As discussed in Sect. 3.1, the relatively high concentrations of $DMAH^+$ and $TMDEAH^+$ over Huaniao Island and the YECS implied that the marine sources contributed substantially to these two aminiums. Accordingly, a spatial variation of aminium concentrations was observed over the YECS during the spring cruise. The concentrations of $DMAH^+$ and $TMDEAH^+$ increased by a fold of 3–5 in the southern ECS (average 24.4 and 40.3 ng m^{-3} for the samples of 7–11 Apr. respectively) compared to the YS and northern ECS (average 7.0 and 8.4 ng m^{-3} for the samples of 27 Mar.–5 Apr. respectively) (Fig. 9). This is consistent with the noticeable difference of Chl-a concentrations between the southern and northern YECS (2.3 folds higher in southern YECS than that in northern YECS, unpublished data). Furthermore, the highest $TMDEAH^+$ and lowest NH_4^+ concentrations observed on 7–11 Apr. corresponded to the air-mass back trajectories originating from the ocean, suggesting that the metabolic activities of surface plankton in the high-productive seas could be a strong source of amines as previously reported (Facchini et al., 2008; Müller et al., 2009; Sorooshian et al., 2009; Hu et al., 2015). Differently, the high concentrations of aminiums observed on 14 Apr. near Qingdao was affected by the air masses transported from eastern China (Fig. 9) and thereby contributed mainly by terrestrial sources.

Fine-mode NH_4NO_3 could decompose during its transport from the land to the ocean, and the released HNO_3 gas would react with dust and sea salt aerosols to form coarse-mode NO_3^-. Therefore, negative correlations were observed between the concentrations of fine-mode NO_3^- and alkaline species (Na^++Ca^{2+}) over the East Asia (Bian et al., 2014; Uno et al., 2017). Since only one dust event was encountered on 12–13 Apr. during the cruise (unpublished data), the coarse-mode NO_3^- in this study should be mostly formed by the heterogeneous reaction with sea salts. Therefore, the importance of terrestrial transport to marine aerosols could be roughly estimated by the percentage of NO_3^- in the fine mode. For aerosols collected on 29–31 Mar., 4–5 Apr., 7–9 Apr. and 9–11 Apr., over 2/3 concentrations of NO_3^- were in the coarse mode (>1.8 μm, Fig. 10a). These samples should be less affected by the terrestrial air masses (referred to category 1) compared to other samples (referred to category 2), and the judgment was consistent with the pointing directions of back trajectories (Fig. S5). Aminiums were negatively correlated with NH_4^+ for Category 1 samples suggesting that aminiums were probably dominated by marine biogenic sources whereas NH_4^+ was influenced by terrestrial transport (Fig. 10b). For Category 2 samples, a positive correlation was found between aminiums and NH_4^+, indicating that terrestrial sources could contribute significantly to aminiums over the YECS in these cases (Fig. 10c).

3.4.3 Source contributions to aminiums over the coastal sea

Huaniao Island is located in the frontline of terrestrial transport to the ECS and influenced by the air masses from the land or ocean depending on the seasonal variation of prevailing winds. Significantly positive correlations were found between the concentrations of aminiums and NH_4^+ in the autumn but not in the summer of 2016 or in late summer of 2017 (Fig. 11). Accordingly, the majority of backward trajectories pointed to the northern China in autumn whereas air masses predominantly originated from the ECS in summer (Fig. 12). Meanwhile, NO_3^- demonstrated a tri-modal distribution with three peaks at 0.18–0.32 μm (condensation mode), 0.56–1.0 μm (droplet mode) and 3.2–5.6 μm (coarse mode) in autumn but only one peak at 3.2–5.6 μm in late summer of 2017 (Fig. S6). These implies that terrestrial transport could be a dominant source for aminiums over the coastal ECS in autumn while marine sources were dominant in late summer. In early summer of 2017, the mass ratios of aminiums to NH_4^+ were significantly lower on 26–28 Jun. than on other days (Fig. S7), corresponding to different origins and properties of the air masses. Removing the data measured on 26–28 Jun., we found a significantly positive correlation

between the concentrations of $DMAH^+$ and NH_4^+ but not between $TMDEAH^+$ and NH_4^+. This suggests that $DMAH^+$ and $TMDEAH^+$ may be predominantly derived from terrestrial and marine sources, respectively.

Good positive correlations were generally found between the concentrations of $TMDEAH^+$ and $DMAH^+$ over Huaniao Island and the YECS, and the slope for autumn samples dominated by terrestrial sources was significantly lower than those influenced primarily by marine air masses (e.g. late summer at Hunaiao Island and spring over the YECS, Fig. 13). The highest slope of $TMDEAH^+$ vs $DMAH^+$ (1.98) occurred in the summer of 2016 which was also mainly affected by marine sources. Therefore, it is speculated that aminiums derived from marine biogenic source might have significantly higher $TMDEAH^+$ to $DMAH^+$ ratios than those from terrestrial sources. Similarly, Hu et al. (2015) observed a significant correlation between the $TMDEAH^+$ and $DMAH^+$ concentrations over the Yellow Sea with the slope of 1.27–2.49. In early summer of 2017, the weak correlation between the $DMAH^+$ and $TMDEAH^+$ and very low slope (0.29) suggested the mixing of terrestrial and marine influence on aminiums over Huaniao Island during that period as discussed above.

The dimethylsulfide (DMS) produced in seawater by the metabolism of plankton will be released into the atmosphere, and SO_2, MSA, SO_4^{2-} and other products can be formed through a series of oxidation reactions(Saltzman et al., 1985; Charlson et al., 1987; Faloona, 2009; Barnes et al., 2006). MSA is often used as a tracer of marine biogenic source to calculate the marine biogenic contribution to nss-SO_4^{2-} (Yang et al., 2009; Yang et al., 2015). Therefore, the mass ratio of MSA to nss-SO_4^{2-} (MSA/nss-SO_4^{2-}) can be used to indicate the contribution of marine sources to aerosol components. A significantly linear relationship was found between aminium/NH_4^+ and MSA/nss-SO_4^{2-} for the samples collected in the autumn of 2016 and summer of 2017 over Huaniao Island (Fig. 14). The value of aminium/NH_4^+ increased with the increasing contribution of marine sources to the aminium. When the marine biogenic source contribution is 0, the corresponding aminium/NH_4^+ values (b in Eq. (3)) represent the average ratios completely contributed by terrestrial sources. By multiplying the ratios by NH_4^+ concentrations, the aminiums contributed by terrestrial sources can be calculated (Eq. (4)). Therefore, the contributions of terrestrial and marine sources to aerosol aminiums can be quantitatively estimated.

$$([aminium]/[NH_4^+])_{terrestrial} = k \times ([MSA]/[nss - SO_4^{2-}])_{terrestrial} + b \quad (3)$$

$$[aminium] = ([aminium]/[NH_4^+])_{terrestrial} \times [NH_4^+] + [aminium]_{marine} \quad (4)$$

where k and b are the slope and intercept of the linear fitting equation of $[aminium]/[NH_4^+]$ and $[MSA]/[nss - SO_4^{2-}]$, respectively (Fig. 14).

Although most of MSA comes from marine sources, the terrestrial sources may also have a certain contribution (Yuan et al., 2004). Therefore, MSA/nss-SO_4^{2-}=0 was not used as the end member value for calculating the terrestrial contribution. In winter, due to the prevailing northwest monsoon and low marine biogenic activities at low temperature, the aerosol components over Huaniao Island were overwhelmingly affected by terrestrial transport. We conducted total suspended particles (TSP) sampling in the winters of both 2014 and 2015 and obtained a total of 41 values of MSA/nss-SO_4^{2-} which were between 0.0010 and 0.0068. The smallest 5 values were considered to represent the situations completely contributed by terrestrial sources, with an average 0.0018±0.0007. Substituting it into the previous fitting equation, the values of $([DMAH^+]/[NH_4^+])_{terrestrial}$ and $([TMDEAH^+]/[NH_4^+])_{terrestria}$ were 0.0062 (0.0044–0.0093) and 0.0028 (0.0008–0.0052), respectively. Then the average contributions of terrestrial and marine sources to the two aminiums in each campaign were calculated and shown in Table 3. It can be seen that the average terrestrial contributions to $DMAH^+$ and $TMDEAH^+$ were both more than 60% in autumn, higher than those in summer. The contributions of marine sources during late summer of 2017 (66.5% for $DMAH^+$ and 82.5% for $TMDEAH^+$) were higher than those in early summer (57.3% for $DMAH^+$ and 79.1% for $TMDEAH^+$), which was consistent with previous speculation. Furthermore, the contribution of marine sources was greater to $TMDEAH^+$ than to $DMAH^+$ in all campaigns, which corresponded to the higher ratio of $TMDEAH^+$/$DMAH^+$ in the samples influenced primarily by marine air masses (Fig. 13). It should be pointed out that although NH_4^+ was mainly derived from the land, marine sources may also had a certain contribution (Altieri et al., 2014; Paulot et al., 2015). This was neglected in our calculation and might lead to the overestimate of terrestrial contributions to aminiums. Besides, the relatively small number of data points used in the fitting (25

326 points) and the treatment of $([aminium]/[NH_4^+])_{terrestrial}$ as a fixed value ignoring its variation would cause uncertainty in
327 the results. Nonetheless, this is the first quantitative estimate of the contributions of terrestrial and marine sources to aerosol
328 aminiums over the coastal ECS, and the method using MSA/nss-SO_4^{2-} as an indicator of marine source is rational and feasible.

329 **4 Conclusion**

330 Amines in the atmosphere play an important role in new particle formation and subsequent particle growth, and studying
331 aerosol aminiums can provide insight into the sources, reaction pathways and environmental effects of amines. An integrated
332 observation was conducted on aerosol aminiums mainly $DMAH^+$, $TMDEAH^+$ and $TEAH^+$ in a coastal city (Shanghai), a nearby
333 island (Huaniao) and the marginal seas (the YECS). All three aminiums exhibited significantly seasonal variation in Shanghai
334 with their highest concentrations in winter, which was consistent with relatively severe air pollution associated with the winter
335 monsoon (continental winds) and the lowest BLH and temperature in this season. Atmospheric oxidizing capacity and
336 relatively humidity may also influence the concentrations of aerosol aminiums to some extent by oxidizing gaseous amines
337 and enhancing the gas-particle partitioning, respectively. By comparing the ocean sites to Shanghai, similar concentrations of
338 $DMAH^+$ and 3-fold higher $TMDEAH^+$ were observed suggesting that these two aminiums may have significant marine sources.
339 Differently, $TEAH^+$ was most abundant aminium in Shanghai but it was below the detection limit over Huaniao Island and the
340 YECS, implying its terrestrial origin.
341 Aminiums influenced substantially by terrestrial transport showed a bimodal distribution with two peaks at 0.18–0.32 μm
342 (condensation mode) and 0.56–1.0 μm (droplet mode), suggesting that the gas-to-particle condensation and cloud processing
343 were primary formation pathways for aerosol aminiums. Nonetheless, aminiums demonstrated a unimodal distribution with a
344 wide peak at 0.18–1.0 μm over the YECS and in late summer of Huaniao Island, and the elevated concentration at 0.32–0.56
345 μm might be related to sea-spray aerosols that either contain primary aminiums or provide surface for heterogeneous reactions
346 to form secondary aminiums. This indicates that aminiums in marine aerosols may undergo different formation pathways from
347 those on land.
348 We firstly distinguished the contributions of terrestrial and marine sources to aerosol aminiums by taking the mass ratio of
349 MSA to nss-SO_4^{2-} as an indicator of marine biogenic sources. In the autumn of 2016, the contributions of terrestrial sources to
350 aminums over Huaniao Island were estimated to be more than 60%. In contrast, marine biogenic sources dominated aminium
351 concentrations especially for $TMDEAH^+$ (~80%) in the summer of 2017. The proposed quantitative estimates may be helpful
352 for simulating the source emissions of amines in atmospheric chemistry models in the coastal area.
353
354 *Data availability.* Data are available from the corresponding author on request (yingchen@fudan.edu.cn).
355
356 *Author contribution.* SZ, YC and CD conceived the study. SZ, YC and CD wrote the paper. SZ, HL, and JX collected the
357 samples. SZ, TY and JX performed the measurement. All have contributed to review of the manuscript.
358
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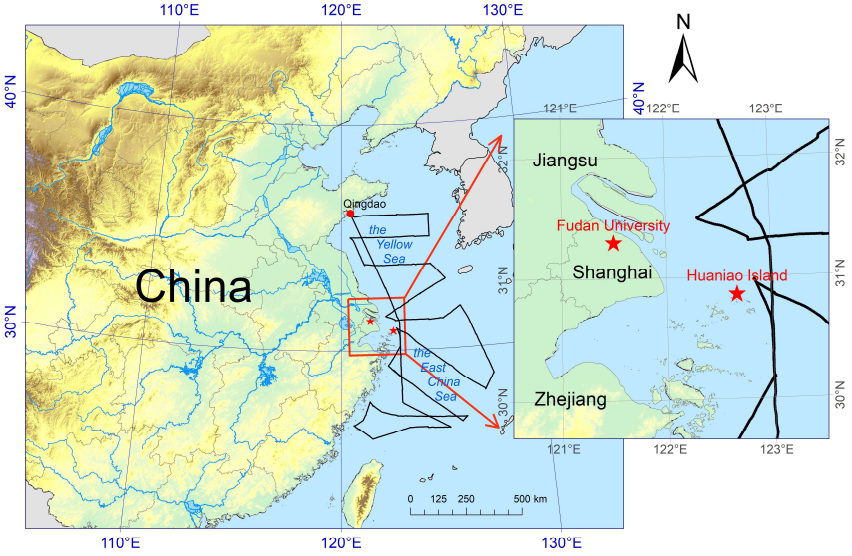

Figure 1. Map of sampling sites and area. The red stars represent the locations of Shanghai (Fudan University) and Huaniao Island, and the black line in the marginal seas represents the cruise track in the spring of 2017.

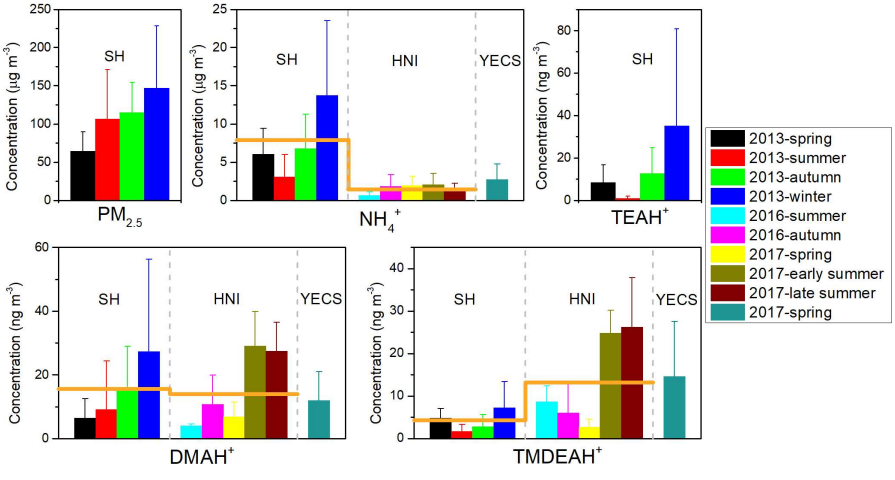

Figure 2. The mass concentrations of $PM_{2.5}$, fine-particle NH_4^+ and three aminiums ($TEAH^+$, $DMAH^+$ and $TMDEAH^+$) in different campaigns in Shanghai (SH), Huaniao Island (HNI) and the Yellow and East China seas (YECS). The columns and error bars represent average concentrations and standard deviations, respectively. The orange horizontal lines represent the annual average concentrations of aminiums in SH and HNI.

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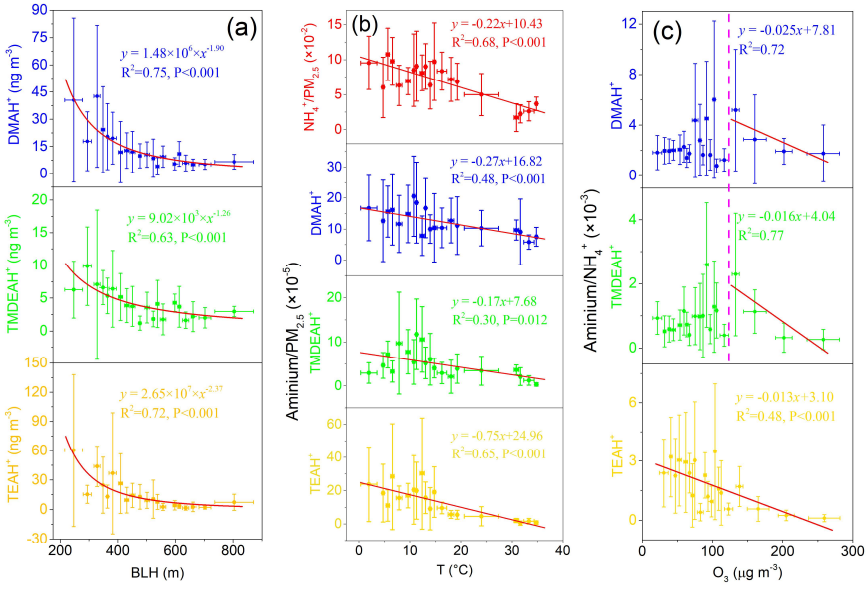

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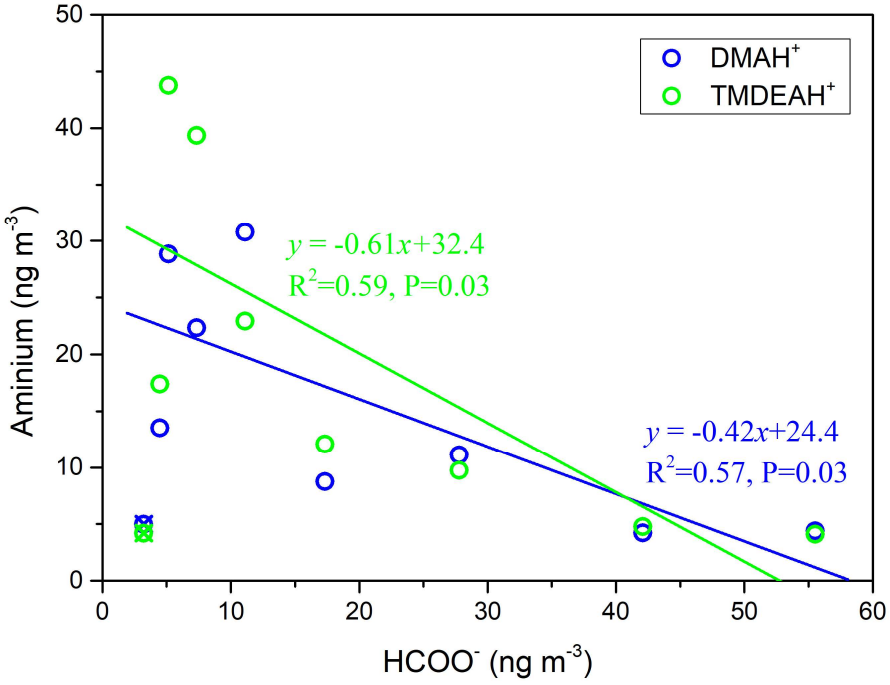
589 **Figure 3. (a)** Relationships between concentrations of aminiums and boundary layer height (BLH). **(b)** Relationships between mass ratios
590 of aminiums and NH_4^+ to $PM_{2.5}$ and temperature. **(c)** Relationships between mass ratios of aminiums to NH_4^+ and O_3 concentrations.



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592 **Figure 4.** Correlations between concentrations of aminiums and $HCOO^-$ over the Yellow and East China seas (YECS) in the spring of 2017.

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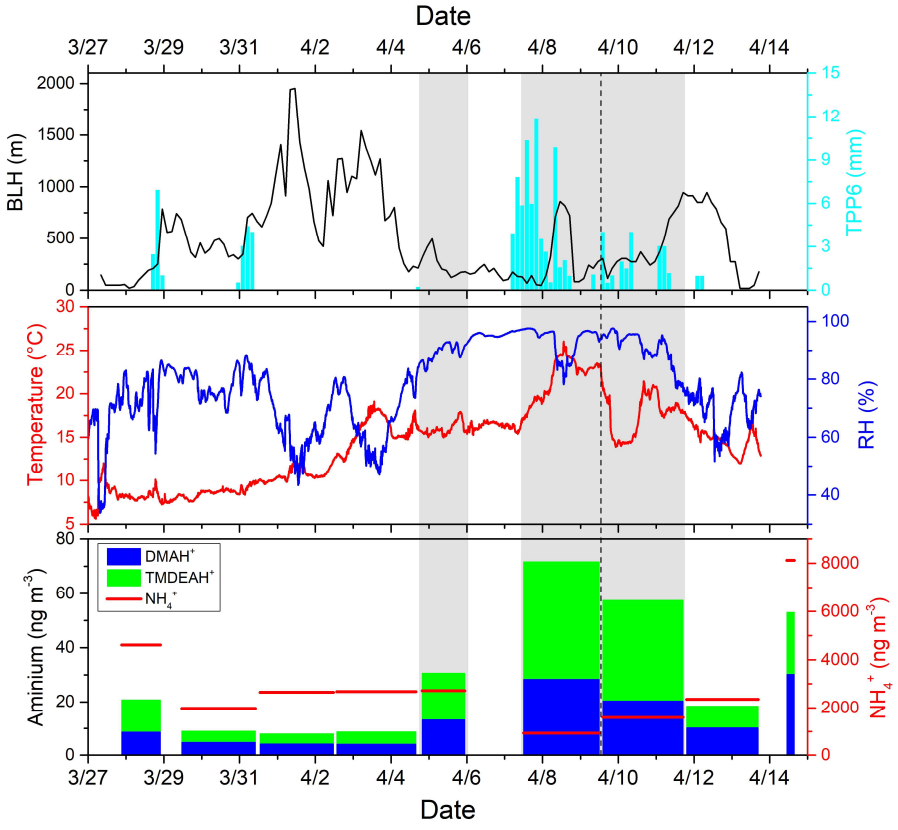

Figure 5. Time series of meteorological parameters and the concentrations of aminiums and NH_4^+ during the cruise of 2017. The time range spanned by the column of each aminium concentration corresponds to the sampling time.

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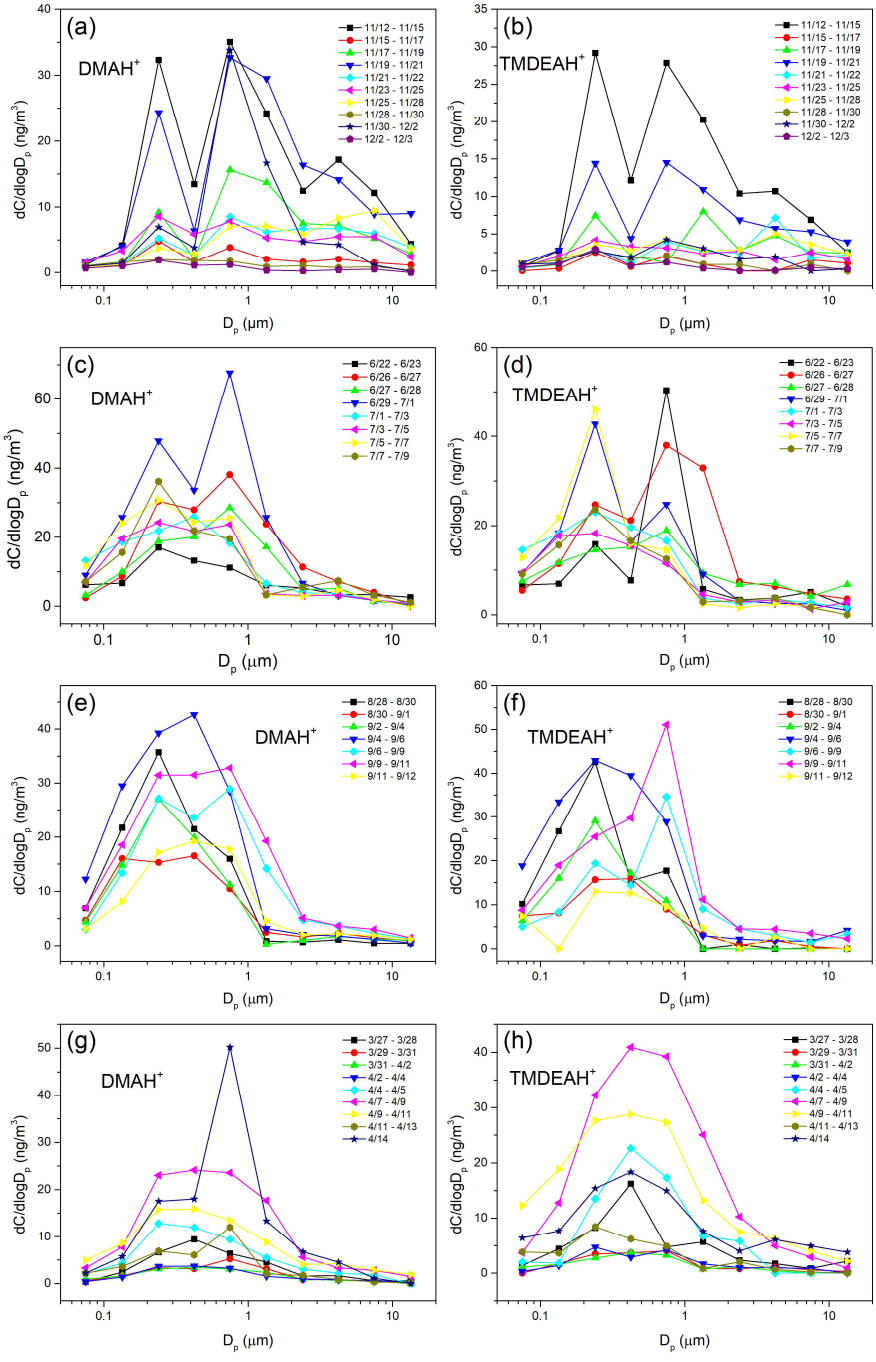

Figure 6. Size distributions of aminiums during different campaigns. **(a-b):** in the autumn of 2016 at Huaniao Island, **(c-d):** in early summer of 2017 at Huaniao Island, **(e-f):** in late summer of 2017 at Huaniao Island, **(g-h):** in 2017 spring cruise over the Yellow and East China seas.

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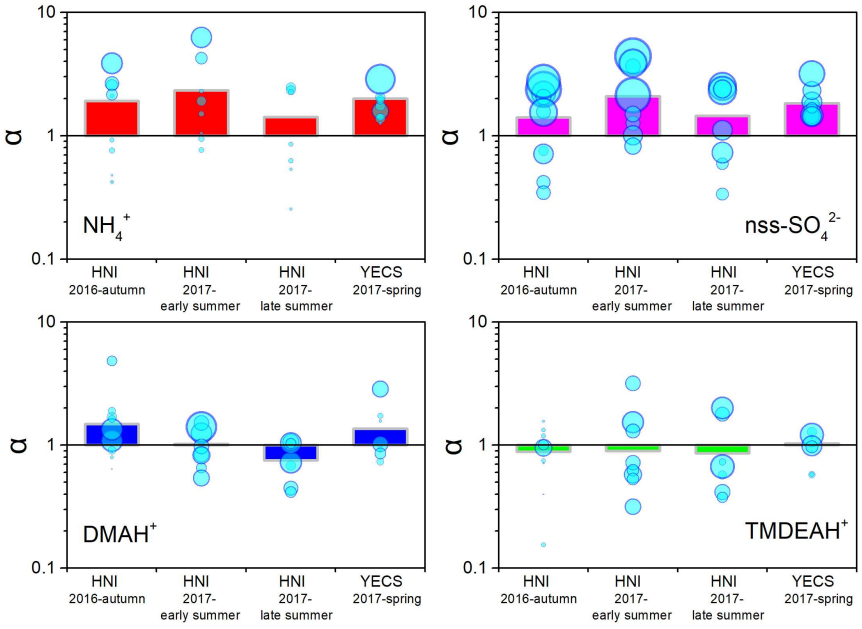

Figure 7. The α values of NH_4^+, nss-SO_4^{2-} and aminiums in different campaigns. The diameter of the circle is proportional to the concentration and the column is the average value of α for each campaign. It should be noted that the bottom of column is the line of α=1.

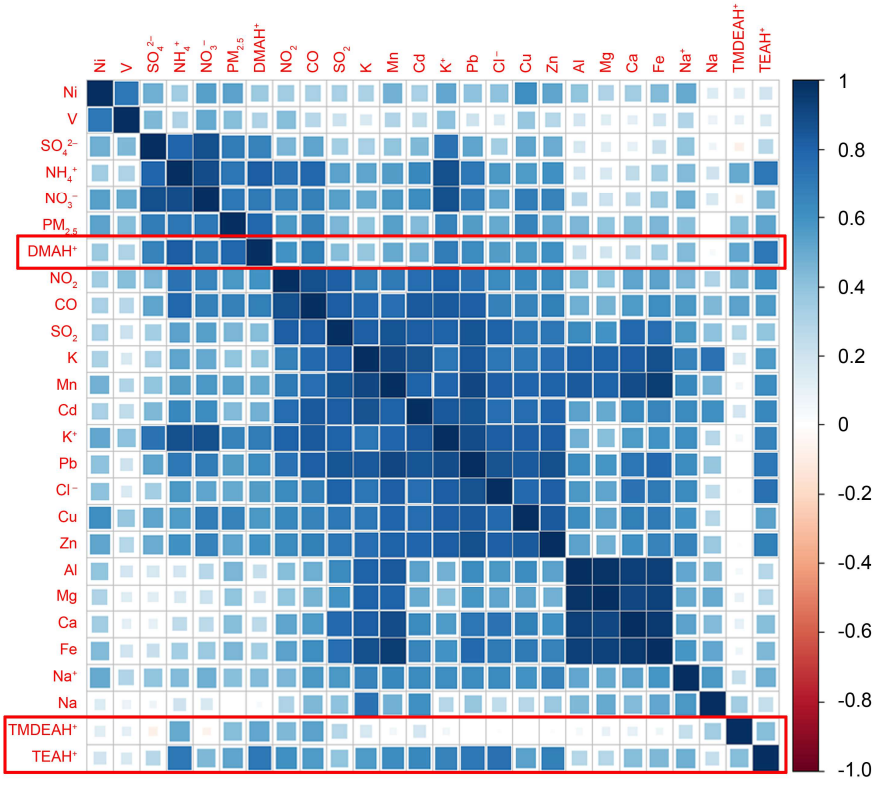

Figure 8. Correlation coefficient matrix among the concentrations of $PM_{2.5}$ components and gaseous pollutants over Shanghai in 2013.

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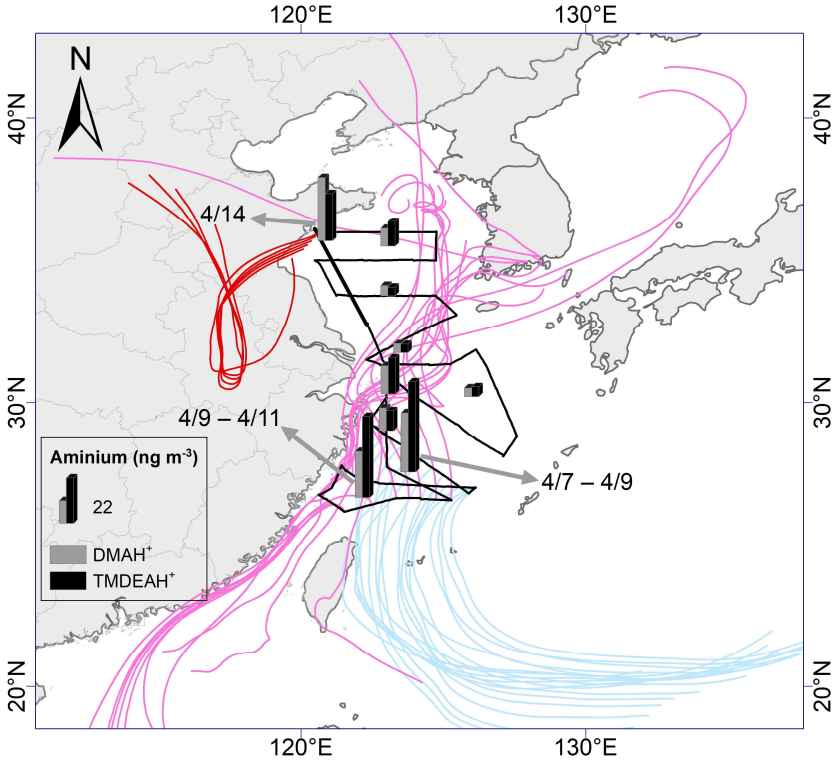

Figure 9. The spatial distribution of aminiums over the YECS in the spring of 2017. The ocean color represents the concentration of chlorophyll a obtained from Kriging interpolation from the observed concentrations. The light blue, pink and red lines represent 72-hour backward trajectories corresponding to sample sets collected on 7–9 Apr., 9–11 Apr. and 14 Apr., respectively.

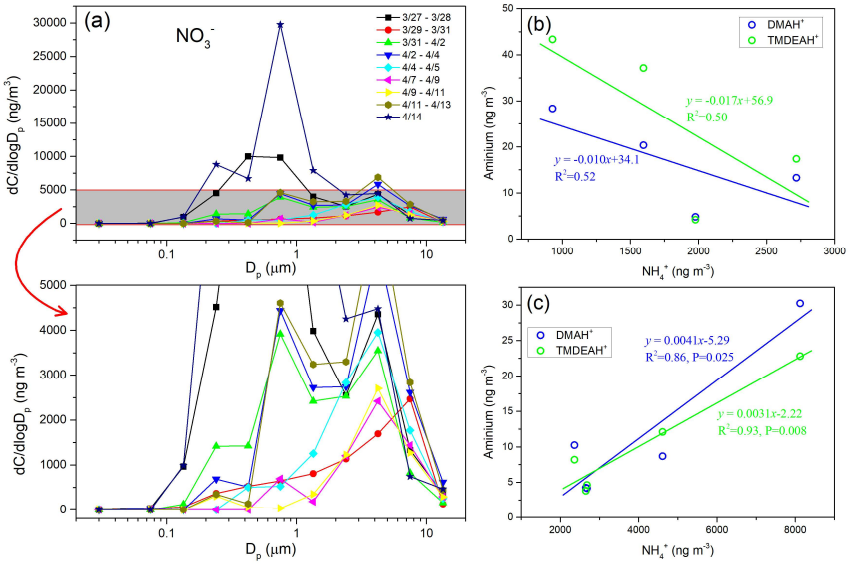

Figure 10. (a) Size distributions of NO_3^- over the YECS in the spring of 2017. **(b)** Correlations between concentrations of aminiums and NH_4^+ for the samples mainly influenced by marine air masses. **(c)** Correlations between concentrations of aminiums and NH_4^+ for the samples predominantly influenced by terrestrial transport.

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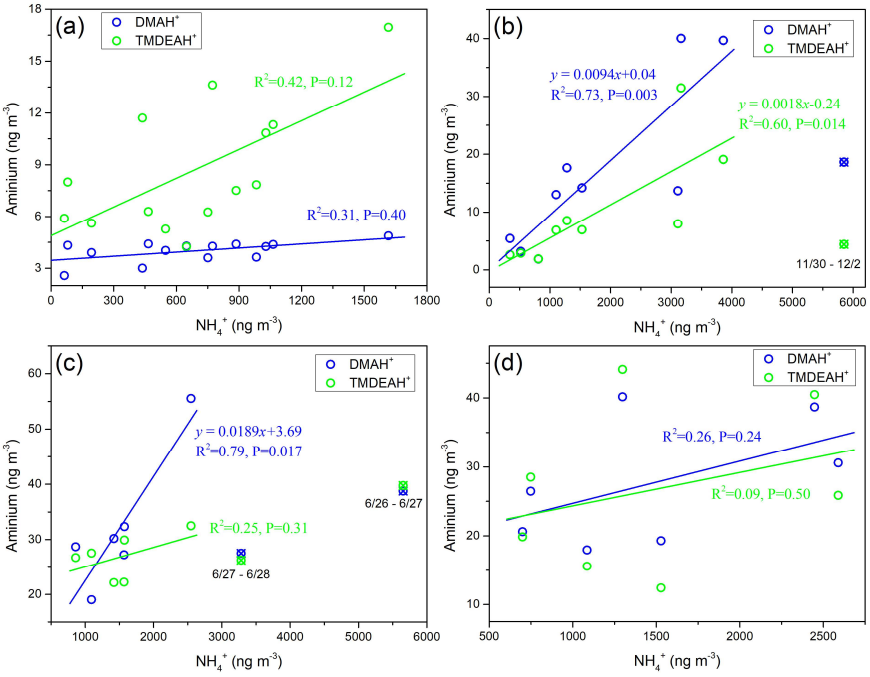


Figure 11. Correlations between aminiums and NH_4^+ concentrations over Huaniao Island for each campaign. **(a):** in the summer of 2016, **(b):** in the autumn of 2016, **(c):** in early summer of 2017, **(d):** in late summer of 2017.

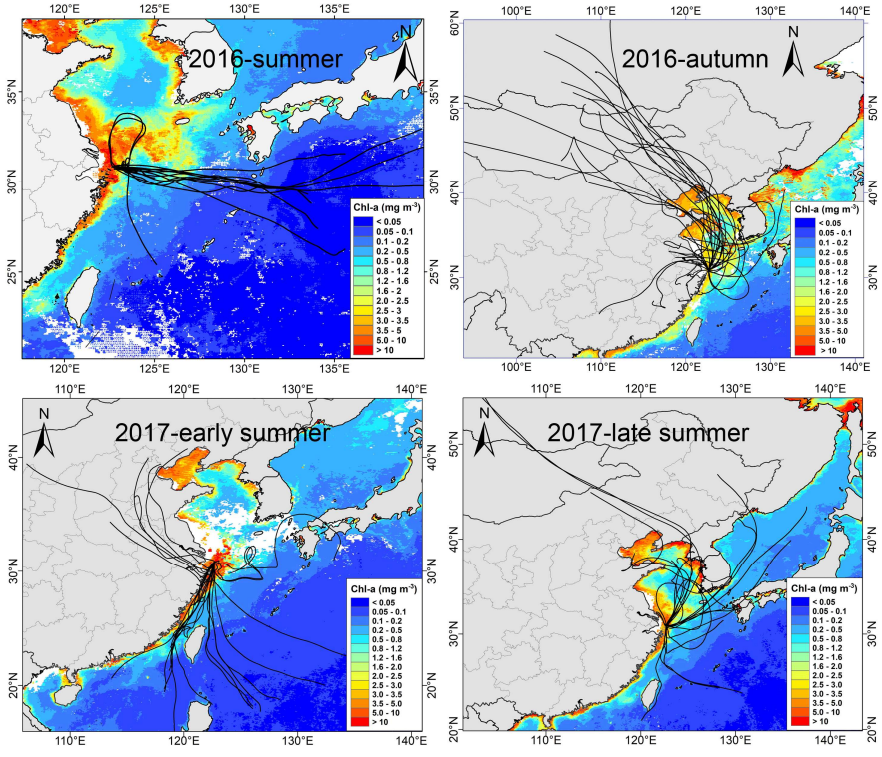
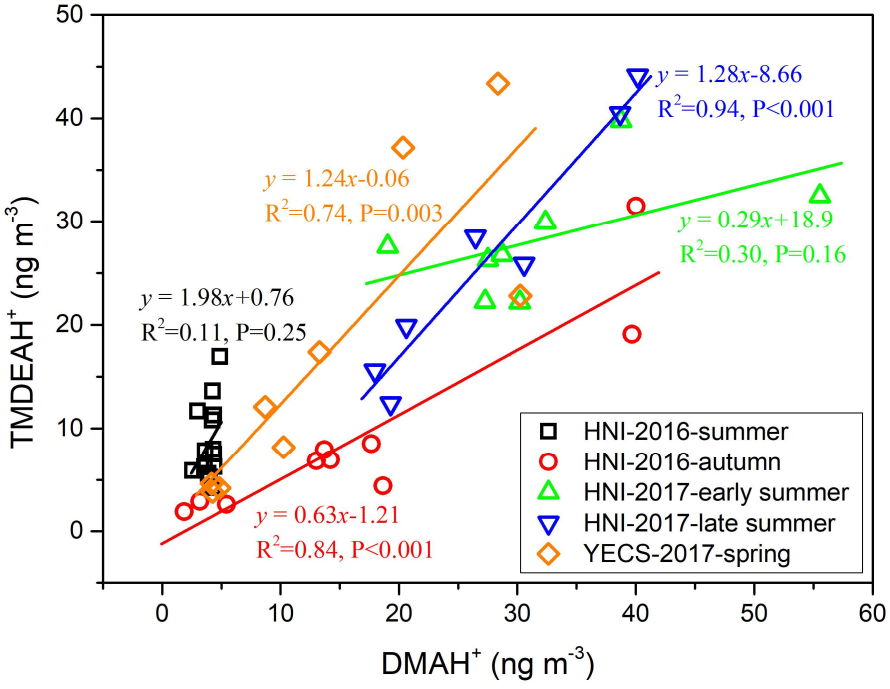


Figure 12. The 72-hour backward trajectories starting from Huaniao Island and the average chlorophyll a concentration retrieved and combined from aqua- and terra-MODIS during the sampling period. Each sample during the summer of 2016 corresponds to one trajectory

618 with a starting time in the middle of sampling period. Each sample set during the autumn of 2016 and the summer of 2017 corresponds to 3
619 trajectories and the starting times are taken at equal intervals in the sampling period.



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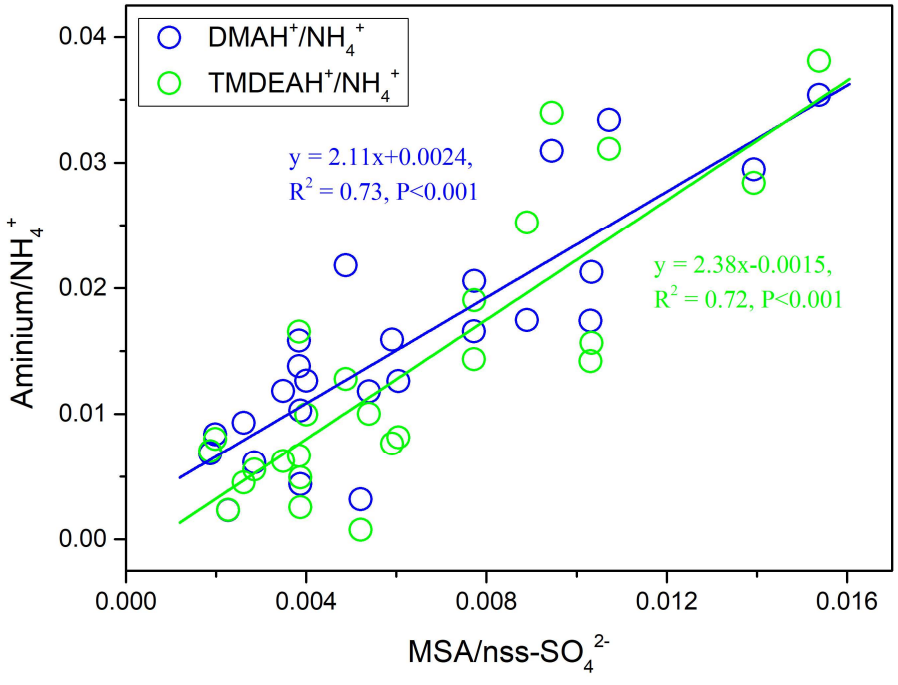
621 **Figure 13.** Correlations between $DMAH^+$ and $TMDEAH^+$ for each campaign over Huaniao Island and the YECS.



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623 **Figure 14.** Correlations between aminium/NH_4^+ and MSA/nss-SO_4^{2-} over Huaniao Island during the autumn in 2016 and the summer in
624 2017.

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626 **Table 1.** Summary of sampling information in different campaigns.

Sampling site	Sampler	Sampling period	Number of samples or sample sets
Fudan University, Shanghai	Medium-flow $PM_{2.5}$ sampler	25 Mar. 2013–26 Apr. 2013 (spring)	29
		16 Jul. 2013–17 Aug. 2013 (summer)	26
		30 Oct. 2013–30 Nov. 2013 (autumn)	29
		1 Dec. 2013–23 Jan. 2014 (winter)	47
Huaniao Island	Medium-flow $PM_{2.5}$ sampler	4 Aug. 2016–18 Aug. 2016 (summer)	14
Huaniao Island	MOUDI	12 Nov. 2016–3 Dec. 2016 (autumn)	9
		11 Mar. 2017–19 Mar. 2017 (spring)	4
		22 Jun. 2017–9 Jul. 2017 (early summer)	8
		27 Aug. 2017–12 Sep. 2017 (late summer)	7
the Yellow Sea and the East China Sea	MOUDI	27 Mar. 2017–14 Apr. 2017 (spring)	9

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29 **Table 2.** The mass concentrations of NH_4^+ and aminiums over Shanghai, Huaniao Island and the YECS compared to other sites reported in literatures. The values below the detection limits are indicated by < DL.

No.	Site	Site type	Sampling period	Particle size	NH_4^+ (μg m^{-3})	Aminium (ng m^{-3})					Reference
						$MMAH^+$	$DMAH^+$	$TMDEAH^+$	$MEAH^+$	$TEAH^+$	
1	Shanghai	urban	Spring (Mar.–Apr. 2013)	$PM_{2.5}$	6.0±3.4		6.4±6.1	4.8±2.3		8.4±8.4	this study
2			Summer (Jul.–Aug. 2013)	$PM_{2.5}$	3.1±2.9		9.1±15.2	1.7±1.6		0.9±1.0	
3			Autumn (Nov. 2013)	$PM_{2.5}$	6.8±4.5		15.5±13.4	2.8±2.9		12.7±12.2	
4			Winter (Dec. 2013–Jan. 2014)	$PM_{2.5}$	13.7±9.8		27.3±29.0	7.3±6.2		35.2±45.6	
5	Shanghai	urban	Jul.–Aug. 2013	$PM_{1.8}$	2.5±1.3	8.9±6.1	15.7±7.9	38.8±17.0	11.5±17.4		(Tao et al., 2016)
6				PM_{10}	2.6±1.3	9.9±6.9	20.1±10.7	47.0±19.9	15.7±26.4		
7	Shanghai	urban	Jan. 2013	$PM_{2.5}$		2.4			0.2		(Huang et al., 2016)
8			Jul.–Aug. 2013	$PM_{2.5}$		3.9			0.3		
9	Yangzhou	urban	Nov. 2015–Apr. 2016	$PM_{2.5}$		4.9±1.9	4.3±2.4		15.4±8.1		(Shen et al., 2017)
10	Nanjing	urban	Apr.–May 2016	$PM_{2.5}$		7.6	4.2		21.7		
11			Aug. 2014	$PM_{1.8}$		7.2±4.1	18.0±11.7		36.4±18.6		
12	Xi'an	urban	Jul. 2008–Aug. 2009	$PM_{2.5}$		14.4±9.6			3.3±2.4		(Ho et al., 2015)
13	Guangzhou	urban	Sep.–Oct. 2014	$PM_{0.95}$	4.3±1.1	41.8±11.4	14.5±3.2	3.7±0.9	3.2±0.4		(Liu et al., 2017)
14				PM_3	5.1±1.4	50.4±13.7	17.7±3.6	4.8±1.4	4.0±0.5		
15				PM_{10}	5.2±1.4	51.8±13.9	19.0±3.8	5.4±1.6	4.2±0.6		
16	Tampa Bay, Florida	urban	Jul.–Sep. 2005	$PM_{2.5}$	1.4±1.2		31.6±28.3				(Calderón et al., 2007)
17	a traffic site, Milan, Italy	urban	Oct. 2013	TSP	4.2±2.9		90±20			360±20	(Perrone et al., 2016)
18	a limited traffic site, Milan, Italy	urban	Oct. 2013	TSP	4.0±3.0		100±10			420±100	
19	Qingdao	semi-urban	May 2013, Nov.–Dec. 2013, Nov.–Dec. 2015	$PM_{0.056\text{-}10}$			6.3	5.8			(Xie et al., 2018)
20	resort beach site of Qingdao	coastal, rural	Aug. 2016	$PM_{0.056\text{-}10}$			28.5±23.0	9.0±6.6			
21	Egbert, Toronto	agricultural and semi-forested	Oct. 2010	$PM_{2.5}$			0.1±0.2	1±0.6			(VandenBoer et al., 2012)
22	Hyytiälä, southern Finland	boreal forest	Mar. 2015	PM_{10}	0.4±0.1	6.8	1.5	1.1			(Hemmilä et al., 2018)
23			Apr. 2015	PM_{10}	0.1±0.1	2.9	3.1	0.7			
24			Jul. 2015	PM_{10}	0.1±0.1	3.0	8.4±4.9	1.8±1.4	0.4		
25	Nanling, Guangdong	forest	Oct. 2016	$PM_{2.5}$	0.9±0.6	8.8±7.8	2.4±3.2	1.1±1.8			(Liu et al., 2018a)
26			May–Jun. 2017		1.8±1.6	11.9±9.8	5.0±2.2	1.7±1.7			

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No.	Site	Site type	Sampling period	Particle size	NH_4^+ ($\mu g\ m^{-3}$)	Aminium ($ng\ m^{-3}$)					Reference
						$MMAH^+$	$DMAH^+$	$TMDEAH^+$	$MEAH^+$	$TEAH^+$	
27	Huaniao Island	marine	Aug. 2016	$PM_{2.5}$	0.7±0.4		4.0±0.6	8.7±3.7		< DL	this study
28			Nov.–Dec. 2016	$PM_{1.8}$	1.9±1.5		10.7±9.3	6.0±6.8		< DL	
29				PM_{10}	2.1±1.8		15.1±12.4	8.4±8.8		< DL	
30			Mar. 2017	$PM_{1.8}$	2.0±1.2		6.8±4.6	2.7±1.8		< DL	
31				PM_{10}	2.3±1.4		11.4±11.6	3.1±2.2		< DL	
32			Jun.–Jul. 2017	$PM_{1.8}$	2.1±1.4		29.0±10.8	24.8±5.4		< DL	
33				PM_{10}	2.2±1.6		32.2±11.0	27.5±5.7		< DL	
34			Aug.–Sep. 2017	$PM_{1.8}$	1.4±0.7		25.8±8.7	25.0±11.0		< DL	
35				PM_{10}	1.5±0.8		27.4±9.1	26.3±11.6		< DL	
36	the Yellow Sea and the East China Sea	marine	Mar.–Apr. 2017	$PM_{1.8}$	2.8±2.0		11.9±9.0	14.6±12.9		< DL	
37				PM_{10}	3.0±2.2		13.5±10.1	16.6±14.5		< DL	
38	the Yellow Sea and the northwest Pacific	marine	Apr. 2015	$PM_{0.056\text{-}10}$			12.9±10.6	13.2±13.8			(Xie et al., 2018)
39	the East China Sea	marine	Jan. 2016	$PM_{0.056\text{-}10}$			30.8±9.7	12.0±6.6			
40	the Yellow Sea and the Bohai Sea	marine	Aug. 2015, Jun.–Jul. 2016	$PM_{0.056\text{-}10}$			33.3	19.4			
41	the south Yellow Sea	marine	Nov. 2013	$PM_{0.056\text{-}10}$			18.9±16.6	31.8±19.2			
42	the Yellow Sea and the Bohai Sea	marine	May 2012	PM_{11}			202±170	432±426			(Hu et al., 2015)
43	the south Yellow Sea	marine	Nov. 2012	PM_{10}			13.3±4.6	30.0±12.6			(Yu et al., 2016)
44	the north Yellow Sea and the Bohai Sea	marine	Nov. 2012	PM_{10}			-	15.0±6.6			
45	Arabian Sea	marine	Aug.–Oct. 1994	$PM_{0.9}$	0.04	3.2	2.1	0.3			(Gibb et al., 1999)
46			Nov.–Dec. 1994	$PM_{0.9}$	0.1	3.7	11.1	0.5			
47	Mace Head	marine	Jan.–Dec. 2006	PM_1			4.7±6.0	7.6±9.4			(Facchini et al., 2008)
48	Irish Weat Coast	marine	Jun.–Jul. 2006	PM_1			14.7±14.3	14.3±8.7			
49	the island of São Vicente in Cape Verde	marine	May–Jun., Dec. 2007	$PM_{0.14\text{-}0.42}$	0.1	0.1	0.4	0.2			(Müller et al., 2009)
50	off the Central Coast of California	marine	Jul. 2007	PM_1				22			(Sorooshian et al., 2009)
51	the Eastern Mediterranean	marine	2005–2006	PM_1			9.2±36.8	< DL			(Violaki and Mihalopoulos, 2010)

631 **Table 3.** Calculated terrestrial and marine source contributions to aminiums over Huaniao Island.

Campaign	$DMAH^+$		$TMDEAH^+$	
	Terrestrial contribution (%)	Marine contribution (%)	Terrestrial contribution (%)	Marine contribution (%)
2016-autumn	71.2 (59.6–81.9)	28.8 (18.1–40.4)	61.6 (25.1–87.4)	38.4 (12.6–74.9)
2017-early summer	42.7 (30.5–54.7)	57.3 (45.3–69.5)	20.9 (5.8–39.1)	79.1 (12.6–94.2)
2017-late summer	33.8 (24.2–45.4)	66.2 (54.6–75.8)	17.5 (4.9–32.9)	82.5 (67.1–95.1)

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