Molecular understanding of new-particle formation from alpha-pinene between -50 °C and 25 °C

Mario Simon[1], Lubna Dada[2], Martin Heinritzi[1], Wiebke Scholz[3,4], Dominik Stolzenburg[5], Lukas Fischer[3], Andrea C. Wagner[1,6], Andreas Kürten[1], Birte Rörup[2], Xu-Cheng He[2], João Almeida[7,8], Rima Baalbaki[2], Andrea Baccarini[9], Paulus S. Bauer[5], Lisa Beck[2], Anton Bergen[1], Federico Bianchi[2], Steffen Bräkling[10], Sophia Brilke[5], Lucia Caudillo[1], Dexian Chen[11], Biwu Chu[2], António Dias[7,8], Danielle C. Draper[12], Jonathan Duplissy[2,13], Imad El Haddad[9], Henning Finkenzeller[6], Carla Frege[9], Loic Gonzalez-Carracedo[5], Hamish Gordon[11,13], Manuel Granzin[1], Jani Hakala[2], Victoria Hofbauer[11], Christopher R. Hoyle[9,15], Changhyuk Kim[16,17], Weimeng Kong[17], Houssni Lamkaddam[9], Chuan P. Lee[9], Katrianne Lehtipalo[2,18], Markus Leiminger[3,4], Huajun Mai[17], Hanna E. Manninen[7], Guillaume Marie[1], Ruby Marten[9], Bernhard Mentler[3], Ugo Molteni[9], Leonid Nichman[19,a], Wei Nie[20], Andrea Ojdanic[5], Antti Onnela[7], Eva Partoll[3], Tuukka Petäjä[2], Joschka Pfeifer[1,7], Maxim Philippov[21], Lauriane L. J. Quéléver[2], Ananth Ranjithkumar[14], Matti Rissanen[2,22], Simon Schallhart[2,18], Siegfried Schobesberger[23], Simone Schuchmann[7], Jiali Shen[2], Mikko Sipilä[2], Gerhard Steiner[3,b], Yuri Stozhkov[21], Christian Tauber[5], Yee J. Tham[2], António R. Tomé[24], Miguel Vazquez-Pufleau[5], Alexander Vogel[1,7], Robert Wagner[2], Mingyi Wang[11], Dongyu S. Wang[9], Yonghong Wang[2], Stefan K. Weber[7], Yusheng Wu[2], Mao Xiao[7], Chao Yan[2], Penglin Ye[11,25], Qing Ye[11], Marcel Zauner-Wieczorek[1], Xueqin Zhou[1,9], Urs Baltensperger[9], Josef Dommen[9], Richard C. Flagan[17], Armin Hansel[3,4], Markku Kulmala[2,13,20,26], Rainer Volkamer[6], Paul M. Winkler[5], Douglas R. Worsnop[2,10,25], Neil M. Donahue[11], Jasper Kirkby[1,7], and Joachim Curtius[1]

[1] Institute for Atmospheric and Environmental Sciences, Goethe University Frankfurt, Frankfurt am Main, 60438, Germany.
[2] Institute for Atmospheric and Earth System Research, University of Helsinki, Helsinki, 00014, Finland.
[3] Institute for Ion and Applied Physics, University of Innsbruck, Innsbruck, 6020, Austria.
[4] Ionicon Analytik, Ges.m.b.H., Innsbruck, 6020, Austria.
[5] Faculty of Physics, University of Vienna, Vienna, 1090, Austria.
[6] Department of Chemistry & CIRES, University of Colorado Boulder, Boulder, CO, 80309-0215, USA.
[7] CERN, Geneva, 1211, Switzerland.
[8] Faculdade de Ciências, Universidade de Lisboa, Lisboa, 1749-016, Portugal.
[9] Laboratory of Atmospheric Chemistry, Paul Scherrer Institute, PSI, Villigen, 5232, Switzerland
[10] TOFWERK AG, Thun, 3600, Switzerland.
[11] Center for Atmospheric Particle Studies, Carnegie Mellon University, Pittsburgh, PA, USA.
[12] Department of Chemistry, University of California, Irvine, CA, 92697, USA.
[13] Helsinki Institute of Physics, University of Helsinki, Helsinki, 00014, Finland.
[14] School of Earth and Environment, University of Leeds, Leeds, LS2 9JT, UK.
[15] Institute for Atmospheric and Climate Science, Swiss Federal Institute of Technology, Zurich, 8092, Switzerland.
[16] School of Civil and Environmental Engineering, Pusan National University, Busan, 46241, Republic of Korea.
[17] Division of Chemistry and Chemical Engineering, California Institute of Technology, Pasadena, CA, 91125, USA.
[18] Finnish Meteorological Institute, Helsinki, 00560, Finland.
[19] School of Earth and Environmental Sciences, University of Manchester, Manchester, M13 9PL, UK.
[20] Joint International Research Laboratory of Atmospheric and Earth System Sciences, School of Atmospheric Sciences, Nanjing University, Nanjing, Jiangsu Province, China.
[21] P. N. Lebedev Physical Institute of the Russian Academy of Sciences, Moscow, 119991, Russia.

[22] Aerosol Physics Laboratory, Physics Unit, Faculty of Engineering and Natural Sciences, Tampere University, Tampere, Finland.

[23] Department of Applied Physics, University of Eastern Finland, Kuopio, 70211, Finland.

45 [24] IDL-Universidade da Beira Interior, R. Marquês de Ávila e Bolama, Covilhã, 6201-001, Portugal.

[25] Aerodyne Research Inc., Billerica, MA, 01821, USA.

[26] Aerosol and Haze Laboratory, Beijing Advanced Innovation Center for Soft Matter Science and Engineering, Beijing University of Chemical Technology, Beijing, China.

50 [a] present address: Flight Research Laboratory, National Research Council of Canada, Ottawa, ON, K1V 9B4, Canada.

[b] present address: Grimm Aerosol Technik Ainring GmbH & Co KG, Ainring, 83404, Germany.

Correspondence to: Mario Simon (simon@iau.uni-frankfurt.de) and Joachim Curtius (curtius@iau.uni-frankfurt.de)

Abstract. Highly-oxygenated organic molecules (HOMs) contribute substantially to the formation and growth of atmospheric aerosol particles, which affect air quality, human health and Earth's climate. HOMs are formed by rapid, gas-phase
55 autoxidation of volatile organic compounds (VOCs) such as α-pinene, the most abundant monoterpene in the atmosphere. Due to their abundance and low volatility, HOMs can play an important role for new-particle formation (NPF) and the early growth of atmospheric aerosols, even without any further assistance of other low-volatility compounds such as sulfuric acid. Both the autoxidation reaction forming HOMs and their new-particle formation rates are expected to be strongly dependent on temperature. However, experimental data on both effects are limited. Dedicated experiments were performed at the CLOUD
60 (Cosmics Leaving OUtdoor Droplets) chamber at CERN to address this question. In this study, we show that a decrease in temperature (from +25 to -50 °C) results in a reduced HOM yield and reduced oxidation state of the products, whereas the new-particle formation rates ($J_{1.7nm}$) increase substantially. Measurements with two different chemical ionization mass spectrometers (using nitrate and protonated water as reagent ion, respectively) provide the molecular composition of the gaseous oxidation products and a 2-dimensional volatility basis set model (2D-VBS) provides their volatility distribution. The
65 HOM yield decreases with temperature from 6.2 % at 25 °C to 0.7 % at -50 °C. However, there is a strong reduction of the saturation vapor pressure of each oxidation state as the temperature is reduced. Overall, the reduction in volatility with temperature leads to an increase in the nucleation rates by up to three orders of magnitude at -50 °C compared with 25 °C. In addition, the enhancement of the nucleation rates by ions decreases with decreasing temperature, since the neutral molecular clusters have increased stability against evaporation. The resulting data quantify how the interplay between the temperature-
70 dependent oxidation pathways and the associated vapor pressures affect biogenic new-particle formation at the molecular level. Our measurements therefore improve our understanding of pure biogenic new-particle formation for a wide range of tropospheric temperatures and precursor concentrations.

1 Introduction

75 Atmospheric aerosol particles play a key role in the regulation of climate by influencing the Earth's radiative energy balance (Stocker et al., 2013). In order to affect the solar radiation budget by acting as cloud condensation nuclei (CCN), newly-formed particles have to reach a size of 50 to 100 nm (Dusek et al., 2006), i.e., they need to grow fast enough to avoid coagulation-scavenging by pre-existing particles. Furthermore, fine airborne particles affect the air quality, are responsible for most air-pollution-related diseases, and cause millions of premature deaths worldwide (WHO, 2016).

80 Around half of the global CCN originate from nucleation of organic or inorganic atmospheric vapours (Spracklen et al., 2008; Merikanto et al., 2009; Kulmala et al., 2013; Gordon et al., 2017). New-particle formation is observed in many environments and under various conditions around the globe, from remote locations such as forested areas or marine/coastal regions to polluted urban areas; from warm environments, such as the tropics, to cold polar and alpine regions; and from Earth's surface to the free troposphere (Kulmala et al., 2004; Kerminen et al., 2018). Gaseous sulfuric acid (Ball et al., 1999; Kuang et al.,
85 2008), ammonia (Kirkby et al., 2011; Kürten et al., 2016), amines (Kurtén et al., 2008; Almeida et al., 2013; Kürten et al., 2014), iodine (O'Dowd et al., 2002; Sipilä et al., 2016) and biogenic volatile organic compounds (BVOCs)(Donahue et al., 2013; Riccobono et al., 2014; Kirkby et al., 2016; Bianchi et al., 2016) have been identified as key vapours involved in atmospheric new-particle formation. The relative importance of each these precursors, however, depends on the particular ambient conditions. The chemical composition of the newly-formed particles is also widely influenced by volatile organic
90 compounds (VOCs), which undergo atmospheric reactions to form secondary organic aerosols (SOA) (Jimenez et al., 2009; Hallquist et al., 2009; Riipinen et al., 2012).

BVOCs emitted by vegetation comprise the dominant fraction of all VOCs, with an estimated global emission rate of 760 Tg (C) per year. Monoterpenes contribute approximately 11 % of all BVOC emissions (Sindelarova et al., 2014). The dominant monoterpene from vegetation (e.g., coniferous trees) is α-pinene, accounting for ~34 % of the total global monoterpene
95 emissions. Most of its oxidation products lead to oxidized volatile organic compounds (OVOCs) with a low degree of oxygenation; they are characterized as intermediate-volatility or semi-volatile organic compounds (IVOC, $300 < C^*(T) < 3 \cdot 10^6 \, \mu g \, m^{-3}$; and SVOC, $0.3 < C^*(T) < 300 \, \mu g \, m^{-3}$, where $C^*(T)$ is the effective saturation concentration, respectively). However, α-pinene has an endocyclic carbon double bond; oxidation of this functionality by ozone can initiate a rapid oxidation process, known as autoxidation (Crounse et al., 2013). Autoxidation proceeds by repeated intramolecular
100 hydrogen shifts (H-shift) of weakly bound hydrogen atoms within peroxy radicals (RO_2). Each H-shift is followed by rapid addition of molecular oxygen (O_2) to form multifunctional peroxy radicals with a high degree of oxygenation, while preserving the radical functionality. Under low NO conditions (Berndt et al., 2018a), these radicals terminate to yield organic products with a high degree of oxygenation and therefore low volatility. Although multifunctional RO_2 radicals formed in autoxidation process are an important intermediate class of compounds in atmospheric chemistry (Rissanen et al., 2015), knowledge about
105 their complex formation mechanisms and kinetics remains far from complete (Ehn et al., 2017).

The autoxidation pathway leads to highly-oxygenated organic molecules (HOMs) with molar yields of several percent (7 % at 20 °C (Ehn et al., 2014); 3.2 % at 5 °C (Kirkby et al., 2016)). This class of oxidation products spans a wide range of volatilities from low-volatility and extremely low volatility towards ultra-low volatility organic compounds (LVOC, $3 \cdot 10^{-5} < C^*(\mathrm{T}) < 0.3 \, \mu\mathrm{g} \, \mathrm{m}^{-3}$; ELVOC, $3 \cdot 10^{-9} < C^*(\mathrm{T}) < 3 \cdot 10^{-5} \, \mu\mathrm{g} \, \mathrm{m}^{-3}$; and ULVOC, $C^*(\mathrm{T}) < 3 \cdot 10^{-9} \, \mu\mathrm{g} \, \mathrm{m}^{-3}$, respectively). While

110 the LVOC and ELVOC classes mainly contribute to the growth of embryonic clusters in the atmosphere, the new class ULVOC refers to molecules with sufficiently low saturation vapor pressure that enables them to reach supersaturation and drive pure biogenic nucleation (Donahue et al., 2012; Bianchi et al., 2019; Schervish and Donahue, 2019).

The fate of the α-pinene peroxy radicals (e.g., $C_{10}H_{15}O_{4,6,8,10}$) is mainly influenced by the presence of nitrogen oxides (NO$_x$), hydroxyl radicals (HO$_x$), or peroxy radicals (RO$_2$). Rapid bimolecular reactions terminate the autoxidation chain by forming

115 closed-shell products and consequently affect the chemical composition of the oxidation products and the molar yield of HOMs (Presto et al., 2005; Ng et al., 2007; Ehn et al., 2014; Berndt et al., 2015; Rissanen, 2018). The reactions with NO and HO$_x$ mainly forms semi-volatile and low-volatility organic compounds, which are important for the growth of particles with sizes above a few nanometers (Donahue et al., 2013; Wildt et al., 2014). Since NO concentrations are usually low in areas where BVOC emissions predominate, the loss of RO$_2$ radicals in bimolecular reactions with NO can generally be neglected. In

120 contrast, the RO$_2$ cross-reaction can form higher-molecular-weight accretion products (ROOR) (Donahue et al., 2011; Berndt et al., 2018b; Valiev et al., 2019). As shown by Tröstl et al. (2016) and Lehtipalo et al. (2018), these gaseous dimeric compounds have the ability to condense irreversibly onto atmospheric particles, or even to contribute to the early-stage growth of molecular clusters, since they cover a wide range of volatility from low to ultra-low vapor pressure. Furthermore, they are also potentially important for new-particle formation, especially in environments dominated by biogenic precursors, e.g.,

125 boreal forests (Mohr et al., 2017; Bianchi et al., 2017).

The bimolecular termination reactions have little or no energy barrier. Their rates are therefore only weakly affected by temperature. In contrast, quantum chemical calculations suggest that the intramolecular isomerization through H-shift within the peroxy radicals has a high activation barrier of 84 kJ/mol or more (Rissanen et al., 2014; Kurtén et al., 2015; Schervish and Donahue, 2019). This results in a strong temperature dependence of the autoxidation, which slows down the oxygenation

130 (HOM yield) at lower temperatures. Consequently, the chemical composition of the initial clusters that form from monoterpene oxidation changes at colder temperatures. This was shown in Frege et al. (2018) for ion-induced nucleation of pure HOM particles. Further, chamber studies showed that not only does the SOA formation rate of monoterpene oxidation have a strong temperature dependence but also the final HOM distribution is affected by the autoxidation rate (Saathoff et al., 2009; Kristensen et al., 2017; Quéléver et al., 2019). Additionally, a recent model study by Schervish and Donahue (2019) showed

135 that the first H-shift reaction of the peroxy-radical isomerization is the rate-limiting step of total HOM formation. Stolzenburg et al. (2018) showed that, despite of the reduction in HOM yield, there was no effect on the growth rate of new particles at the lower temperatures. It was shown that the steep exponential temperature dependence in the saturation vapor pressure, as described by the Clausius-Clapeyron relation, counters the reduction of the oxidation state in terms of their volatility distribution. Recent measurements of particle composition by Ye et al. (2019) showed that this leads to sufficient condensation

of even the low- and moderately-oxygenated organic products at low temperature. Therefore, the volatility of the oxidation products is relevant to characterize their ability to condense and participate in new-particle formation. Thus, the VBS model is a suitable tool to track the volatility change of the oxidation of volatile organic compounds with temperature.

Model simulations suggest that highly-oxygenated organic molecules have a pronounced effect on new-particle formation (NPF) on a global scale, especially in pristine environments dominated by biogenic precursors such as the tropical rain forests or at high altitudes as well as in the preindustrial atmosphere (Gordon et al., 2017). Furthermore, recent observations support this conclusion, suggesting that oxidation products of BVOCs have a major impact on the formation of CCN, especially at high altitudes in the tropical convective regions (Williamson et al., 2019). However, the lack of knowledge about the mechanisms and the accurate representation of NPF from BVOCs for different environmental conditions, especially their temperature dependence, remains a great challenge for atmospheric chemistry and climate models.

In the current study, we present a comprehensive investigation of the effect of ambient tropospheric temperature on the molecular composition of α-pinene oxidation products and new-particle formation rates. The experiments were conducted at the CLOUD (Cosmic Leaving OUtdoor Droplets) chamber at CERN (Geneva, Switzerland), using atmospherically-relevant concentrations of α-pinene and ozone. To study pure biogenic nucleation, the addition of other trace gases was avoided in this study. Going beyond the results of Stolzenburg et al. (2018), this study focuses on new-particle formation over a wide range of tropospheric temperatures from ground level (25 °C) to the upper-free troposphere (-50 °C).

2 Methods

2.1 The CLOUD experiment

The CERN CLOUD chamber is a 26.1 m³ electro-polished stainless-steel vessel for the study of new-particle formation under atmospherically relevant conditions. The use of boiled-off nitrogen and oxygen from ultra-clean cryogenic liquids in a ratio of 79:21 minimizes the levels of contaminants (e.g., SO_2, NH_3, NO_x or volatile organics) inside the chamber. CLOUD is operated at a slight overpressure (5 hPa) to avoid contamination at any time, especially when instruments are being connected or disconnected. The relative humidity is adjusted with a temperature-controlled Nafion humidifier using ultra-pure Millipore water. Ozone and other trace gases are introduced by individual gas lines; gas dilution stages are applied when necessary to achieve the targeted mixing ratios.

To add monoterpene, dry nitrogen is passed through a temperature-controlled evaporator containing liquid α-pinene (Sigma-Aldrich, 98%). Efficient uniform mixing of the gases and ions in the chamber is ensured by two magnetically coupled fans located at the bottom and top of the vessel. The characteristic wall loss rates of condensable gases can be adjusted by variation of the fan speed.

The ion concentration in the chamber can be regulated to values that are representative of the full range of tropospheric and stratospheric conditions by the controlled irradiation with a 3.5 GeV/c π^+ beam from the CERN Proton Synchrotron. This simulates the ionizing muon irradiation in the upper troposphere and stratosphere. Furthermore, as the chamber is continuously

exposed to galactic cosmic rays, a 20 kV m^{-1} electrical high voltage clearing field (HVCF) can be imposed by energizing two electrode-grids located at the top and bottom of the chamber, removing all ions within seconds. Thus, the CLOUD chamber enables investigation of new-particle formation under ion-free conditions as well as ion concentrations that are found throughout the troposphere.

Photochemical processes, such as the photo-dissociation of ozone to produce OH radicals, can be controlled by homogeneous illumination with UV light of adjustable intensity. The light from four 200W Hg-Xe UV-lamps (UVH LC8, Hamamatsu Photonics K.K., Japan) is guided by a fiber-optic system into the chamber to avoid any heat load from the light sources and to establish near-homogenous illumination (Kupc et al., 2011).

A thermal housing surrounds the chamber to maintain a high temperature uniformity and to control the chamber temperature in a range from -70 °C to 100 °C with a precision of ±0.1 K. This stability is mandatory as many of the new-particle formation and oxidation processes are highly sensitive to temperature. The temperature inside the chamber is measured with several arrays of thermocouples, while the chamber wall temperature is monitored by a set of calibrated Pt100 sensors (Dias et al., 2017).

Similar to previous CLOUD experiments, state-of-the-art instruments are used to determine the chamber conditions, the concentration of important gas species, and aerosol properties during nucleation and early-growth studies (Kirkby et al., 2016; Lehtipalo et al., 2018; Stolzenburg et al., 2018). All key instruments are placed in the mid-plane of the chamber to ensure sampling from well-mixed conditions inside the chamber. The sampling lines protrude 40 cm into the chamber to avoid sampling close to the walls and to reduce memory effects. Prior to changing to a new chemical system, the chamber and the sampling lines are rinsed with ultrapure water and subsequently heated up to 100 °C to clean the chamber from residual chemicals of previous experiments. Applying high ozone concentrations for several hours during the cleaning helps achieving sub-ppt$_v$ contamination levels of inorganic and <150 ppt$_v$ of total organic compounds (Schnitzhofer et al., 2014). More details about the CLOUD experiment can be found in Kirkby et al. (2011) and Duplissy et al. (2016).

The experiments reported here were performed during the CLOUD10 (Fall 2015), CLOUD12 (Fall 2017) and CLOUD13 (Fall 2018) campaigns. Within these three campaigns, sets of experiments at five different temperatures were performed to study the HOM production and new-particle formation from α-pinene oxidation. α-Pinene was added to the chamber at volume mixing ratios ranging from 100-2000 ppt$_v$, while ozone levels were kept between 30-40 ppb$_v$. OH radicals were mainly formed by the ozonolysis of α-pinene with an 80% yield (Chew and Atkinson, 1996) and also by UV-photolysis of ozone. The relative humidity was commonly held at 40 % in CLOUD10 and CLOUD12 and 80 % in CLOUD13.

Before starting a new-particle formation sequence (run), the CLOUD chamber was cleaned from residual particles and organic compounds by flushing the chamber with clean synthetic air for several hours, while operating the mixing fans at a high speed and periodically activating the HVCF to remove all charged aerosol particles efficiently. The results reported here were obtained without any addition of SO$_2$, NO$_x$ or other trace gases in order to achieve pure biogenic system, to isolate the chemistry of biogenic precursors, and to avoid the interference with other potentially nucleating compounds. Furthermore, no OH radical

205   scavenger was used during the experiments to ensure a faithful simulation of atmospheric conditions. The instruments and methods relevant for the present study are described in the following sections.

2.2 Nitrate CI-APi-TOF

The nitrate CI-APi-TOF uses nitrate anions [$(HNO_3)_n(NO_3^-)$, with n=0-2] as reagent ions which are produced by exposing a nitric acid (HNO_3) enriched sheath gas flow to a corona discharge (Kürten et al., 2011). Based on the free-jet flow design of

210   Eisele and Tanner (1993) the nitrate reagent ions are electrostatically pushed into the sample flow in the center of the ion-molecule reaction drift region without mixing of both gas streams. After a reaction time of 50 ms within the sample flow, the ions and charged clusters enter the atmospheric pressure interface of the mass spectrometer (APi-TOF, Tofwerk AG, Switzerland) where they are focused by two segmented quadrupole units and an ion lens assembly, while the pressure is gradually reduced to around 10^{-6} mbar. In the time-of-flight region, the ions are separated according to their mass-to-charge

215   ratio and counted by a micro-channel plate detector. The data are processed and analyzed in IGOR Pro (WaveMetrics, Inc., USA) using the software package TOFWARE (Version 3.1, Aerodyne Inc., USA).

The chemical ionization with nitrate anions is selective towards strong Lewis acids, like sulfuric acid (H_2SO_4) (Jokinen et al., 2012) or iodic acid (HIO_3) (Sipilä et al., 2016), but also for bases, like dimethylamine ($(CH_3)NH$)) when ion-clusters are being formed, including the nitrate reagent ions (Simon et al., 2016). Highly polar functional groups, like carboxylic acids (COOH),

220   hydroperoxides (R-O-OH) and peroxy acids (R(O)-O-OH), which are the most abundant functional groups in HOMs, can also be detected (Hyttinen et al., 2015). While strong acids are mostly detected as deprotonated anions (e.g., HSO_4^-), HOMs are charged mainly through ion-adduct formation [$HOM_i \cdot NO_3^-$]. Here, the index i denotes a specific HOM molecule (with a specific exact mass). The concentration of the sample is achieved by normalization of the product ion count rates (cps) with the intensity of the reagent ions (cps) expressed by the following Eq. (1):

225

$$[HOM_i] = C \cdot TE_i \cdot SL_{HOM_i} \cdot \ln\left(1 + \frac{[HOM_i \cdot NO_3^-]}{\sum_{j=0}^{2}[(HNO_3)_j \cdot NO_3^-]}\right). \qquad (1)$$

Three different correction factors are considered to obtain a concentration from the raw count rate. First, a general calibration coefficient, C, of the mass spectrometer is applied, which is determined from a calibration using sulfuric acid as described in Kürten et al. (2012). Here, we assume that all HOMs with an oxygen to carbon ratio (O/C) of ≥ 0.6 have a collision-limited charging efficiency when reacting with the nitrate ions similar to sulfuric acid. In addition, we assume that the charging

230   efficiency of the nitrate CI-APi-TOF technique does not change significantly with temperature or humidity (Viggiano et al., 1997). The estimated detection limit of the instrument for sulfuric acid is about $5 \cdot 10^4$ molecules cm^{-3}, however, due to a better signal-to-noise ratio at higher mass-to-charge ratios, some HOM_i can even be quantified at lower concentrations. Second, the mass dependent transmission efficiency TE_i of the instrument is considered by depleting the reagent ions by various perfluorinated acids according to the method described by Heinritzi et al. (2016) in a separate characterization experiment at

235   the beginning and end of the campaign. Third, a temperature-dependent sampling line loss correction factor, SL_{HOM_i}, is considered. It depends on the sample flow rate, the diffusion coefficient of the target molecule and the length of the sampling line. We assume laminar flow diffusional loss in the 120 cm sampling line. To reduce wall losses we applied a core-sampling technique as described by Knopf et al. (2015) and Fu et al. (2019). A fraction of 8.5 standard liters per minute (slm) of the total flow in the inlet line (40 slm) is sampled from its center. This set-up minimizes the section length to 30 cm that transports

240   the sample to the instrument at the smaller flow rate, reducing the sampling loss rate of HOMs to less than 30 %.

As the molecules detected by the nitrate CI-APi-TOF have typically very low saturation vapor pressures, we assume that they are irreversibly lost upon contact with a surface. The diffusion coefficients D_i for each HOM_i are approximated with the expression $D_i[cm^2 \, s^{-1}] = 0.31 \cdot M_i^{-1/3}$, where M_i [g/mol] is the mass of the molecule. The wall loss rate inside the chamber at each temperature is determined from the expression:

$$k_{wall}(T) = C_{wall}(T) \cdot \sqrt{D_i} \, , \tag{2}$$

245   where C_{wall} is an empirical parameter. C_{wall} is derived from dedicated sulfuric acid decay experiments at all relevant temperatures and ranges between 0.0071 and 0.0077 $cm^{-1}s^{-0.5}$ for - 50 °C to + 25 °C. For these experiments the measured wall loss rate and the diffusivity of sulfuric acid (0.078 cm^2s^{-1} at 298 K at a relative humidity of 40 %), is scaled to the chamber temperature by the parametrization $(T/298K)^{1.75}$ (Hanson and Eisele, 2000). The sampling line and the sheath flow of the ion source are thermally insulated and are operated at the same temperature as the CLOUD chamber to avoid evaporation

250   or condensation due to changes in the thermal conditions during the measurements. Note that ions formed in the CLOUD chamber, e.g., from GCRs or from the CERN pion beam, are removed by an electrostatic filter in the nitrate CI-APi-TOF inlet. Therefore, ions from the chamber do not interfere with the CI detection scheme. Finally, the sample ion signals are background corrected for a pure nitrogen sample without any VOC addition to the chamber.

2.3 PTR3-TOF

255   The PTR3-TOF-MS, or PTR3 for short, described in Breitenlechner et al. (2017) uses proton transfer or ligand switching reactions from hydronium water clusters to ionize the majority of organic compounds, specifically those of which have proton affinities larger than that of the water clusters. $H_3O(H_2O)_n^+$ primary ions, produced in a corona discharge from humidified nitrogen, are transferred through a source drift region into the tripole where the ion molecule reactions take place. A core flow of typically 2 slm drawn from the laminar sample gas (10 slm) enters the tripole reaction region through a critical orifice. A

260   pressure controller maintains a constant pressure of typically 70-80 hPa in the reaction region.
By applying a tunable radio frequency signal on the tripole rods, it is possible to adjust the collision energy between ions and sample gas molecules. Elevated collision energies supress cluster ion formation of both primary and product ions but could also lead to unwanted fragmentation of certain product ions. Low collision energies on the other hand increase unwanted clustering of ions with water molecules and decrease the ionization efficiency for molecules with a proton affinity close to that

265   of water. During CLOUD experiments we adjusted the collision energy to E/N values (E being the electric field strength and N the sample gas number density) of 62-72 Td (1 Townsend equals 10^{-17} V cm^2) by using an RF-frequency of 10 MHz and an RF-amplitude of 800-900 V$_{pp}$ at a pressure of 75-77 hPa. With these settings even volatile organic compounds are detected and humidity effects are minor. Primary and product ions were analyzed with a Long-TOF (LTOF, Tofwerk AG, Switzerland). All data were acquired using the TofDAQ recorder by Tofwerk and analyzed with the TOF-Tracer software written by Lukas

270   Fischer running on Julia 0.6 (https://github.com/lukasfischer83/TOF-Tracer).

Precursor molecules are calibrated using a gas standard. More oxidized molecules have typically higher proton affinities; their concentrations are estimated by using the sensitivity of 3-hexanone. Oxidized organic compounds might undergo fragmentation in reactions with $H_3O(H_2O)_n^+$ primary ions, especially when containing hydroperoxide groups (Bernhammer et al., 2017). Therefore, concentrations are lower limit estimates.

275   Furthermore, data are corrected for the duty cycle transmission effects of the TOF and sampling line losses. In Breitenlechner et al. (2017) a correction factor of 5 for the inlet line losses led to good agreement with the nitrate CI-APi-TOF for most highly oxygenated molecules containing more than 5 oxygen atoms in the α-pinene system (Fig. S1). The compounds measured by the PTR3 span several orders of magnitude of volatility, from volatile (VOCs) to extremely low volatility organic compounds (ELVOCs). Therefore, the correction for sampling line losses of less oxidized molecules can only be done by changing the

280   inlet flow rate or the fan speed inside the CLOUD chamber for each inlet temperature and testing the instrument's response for different compounds due to enhanced wall collisions. We then applied a scaled sampling line loss correction factor ranging from 1 (no correction for VOCs, unaffected by changing the number of wall collisions) to 5 (maximum inlet correction for ELVOCs), which decreased during the tests to 20 % (or less) of their value before changed inlet flow or fan speed conditions. Molecules that contain more than 5 oxygen atoms are considered ELVOCs and are automatically corrected by a factor 5 since

285   these compounds are often too close to the detection limit of the PTR3 to get a reasonable response during the tests. Further details about the method is given in Stolzenburg et al. (2018).

2.4 Particle measurements and formation rate determination

The particle number size distributions in the size ranges between 1.2 nm and 1 µm in the chamber were measured by a series of aerosol particle counting instruments. The concentration of the smallest particles was measured with a particle size magnifier

290   (PSM, Airmodus Ltd., Vanhanen et al. (2011)). The PSM was operated in scanning mode for the determination of the particle concentration at different cut-off diameters and for the particle number size distributions between 1 and 3 nm (Wimmer et al., 2013; Lehtipalo et al., 2014; Kürten et al., 2015). Additionally, a butanol CPC (CPC3776, TSI Inc.) with a fixed cut-off diameter of 2.5 nm was used. A DMA-train measured the size distribution of particles in the 1.8 to 8 nm size range with 10 second time resolution; it consists of six differential mobility analyzers (DMAs) with PSM or CPC detectors that are operated

295   in parallel, each measuring a fixed size (Stolzenburg et al., 2017). A commercial scanning nano-SMPS (nSMPS 3982, TSI Inc.) resolved the particle size distribution between 8-63 nm. For larger particles (>50 nm) two additional SMPS systems were used.

The ion concentration and the size distribution of charged clusters and small particles of both polarities were measured using a neutral cluster and air ion spectrometer (NAIS, Airel Ltd., Manninen et al. (2009)). This instrument determines the ion mobility distribution in the range from 0.82 to 45 nm in mobility-equivalent diameter, as well as the total particle size distribution in the size range of 2.5 to 45 nm by charging all sampled aerosols with a periodically activated corona charger. The particle formation rates used in this study were determined using the full particle size distribution following the method presented in Wagner et al. (2017) and Lehtipalo et al. (2018). In practice, the particle formation rates at the cut-off diameter, 1.7 nm, was determined from the time derivatives of the total particle concentrations with diameters larger or equal to 1.7 nm. The formation rates are corrected for the particle losses in the chamber such as dilution, coagulation and wall losses.

2.5 Trace gas and water vapor measurements

Trace gas monitors were used to measure the concentration of ozone (O_3, Thermo Environmental Instruments TEI 49C), sulfur dioxide (SO_2, Thermo Fisher Scientific, Inc. 42i-TLE) and nitrogen oxides (NO, ECO Physics CLD 780TR; NO_2, CE-DOAS (University of Colorado Boulder) and CAPS NO_2 (Aerodyne Research Inc.)). The water vapor concentration in the chamber was monitored with a chilled dew-point mirror (EdgeTech Instruments) and a direct tunable diode laser absorption spectrometer (TDL-hygrometer, Werle et al. (2002)).

2.6 Experimental errors

The overall scale uncertainty for the HOM and oxidation product (OVOC) measurements is +78 %/-68 %. The uncertainty on the formation rates were determined by using the error propagation method of both systematic and statistical uncertainties including those associated with the particle concentration measurement (10 %), as well as their dilution (10 %) and diffusional (20 %) losses. The statistical errors include uncertainty on dN/dt and coagulation sink, which varied from run to run, depending on the stability of the measurement conditions. The reproducibility (run-to-run uncertainty) under identical conditions is about 30 % as described in more detail by Kirkby et al. (2016) and Lehtipalo et al. (2018).

2.7 Volatility Basis Set Model

The ambient temperature and the concentration of the oxidation products significantly determines their saturation vapor pressure. HOM are mainly assigned to the volatility class of LVOC and ELVOC (Bianchi et al., 2019). However, this assignment depends strongly on the temperature. Since the definition of HOM has no direct relation to their physical properties, the volatility classification introduced by Donahue et al. (2011) is used in the present study to discuss the contribution of different HOM and less oxidized products to new-particle formation. In principle, the saturation vapor pressure of an organic molecule is determined by its mass and its functional groups which affects the strength of the interaction with its neighboring molecules, and on the temperature.

The determination of the exact volatility of the oxidation products is challenging because the individual compounds cannot be isolated, as they are highly reactive and fragile species with extremely low saturation vapor pressures. However,

experimentally derived volatilities from desorption thermograms measured with the FIGAERO (Filter Inlet for Gases and

330 AEROsols) inlet show a good agreement with the combination of semi-empirical methods and theoretical model calculations (Lopez-Hilfiker et al., 2014; Schobesberger et al., 2018). This was recently verified in a complementary study of the α-pinene ozonolysis products examined here (Ye et al., 2019), in which the volatility distribution of molecules in the nucleated particles, measured with a FIGAERO inlet over a wide range of temperatures, is in good agreement with those estimated by Stolzenburg et al. (2018).

335 Here we follow the same approach as described in Stolzenburg et al. (2018). We combine the semi-empirical group-contribution methods (SIMPOL, Pankow and Asher (2008)) with the two-dimensional volatility basis set (2D-VBS) introduced by Donahue et al. (2011). It is based on the relationship between a typical molecular composition and its known volatility by parameterizing the saturation vapor pressure of an unknown molecule according to its mass and oxidation state:

$$OS_C = 2 \cdot O/C - H/C. \tag{3}$$

(Donahue et al., 2012; Donahue et al., 2013). Therefore, the volatility can be expressed as the logarithm of the saturation mass

340 concentration, $\log_{10} C_i^*$, from the number of carbon atoms, n_C, and oxygen atoms, n_O, within the specific molecule, i:

$$\log_{10} C_i^* (300K)[\mu g\ m^{-3}] = (n_C^0 - n_C^i) \cdot b_C - n_O^i \cdot (b_O - b_{add}) - 2 \frac{n_C^i \cdot n_O^i}{n_C^i + n_O^i} b_{CO}. \tag{4}$$

Based on Donahue et al. (2011) and a revised version given in Stolzenburg et al. (2018) the parameter $n_C^0 = 25$ represents the baseline carbon backbone for a volatility of 1 μg m⁻³ without the addition of any functional groups. The parameter $b_C = 0.475$ accounts for roughly a half order of magnitude decrease in saturation vapor pressure per carbon atom according to the mass of the molecule while $b_O = 2.3$ considers a more than two orders of magnitude decrease in volatility per oxygen atom assuming

345 an equal proportion of carbonyl (=O) and hydroxyl (-OH) groups in the molecule. The carbon-oxygen nonideal interaction $b_{CO} = -0.3$ is a non-linearity term that adjusts the volatility estimation from organics dominated by carbonyl (=O) and hydroxyl (-OH) groups at low O/C ratios towards HOM, which mainly consist of hydroperoxyl (-OOH) and peroxy acid (-C(O)OOH) groups at high O/C. While the additional oxygen in the -OOH group (log C^* = -2.4) has an almost negligible effect in reducing the saturation vapor pressure compared to the -OH group (log C^* = -2.2), neither covalently bound dimers nor the ability of

350 hydroperoxide and peroxy acid functionalities to form intramolecular hydrogen bonds are included in the non-linear terms b_O and b_{CO} (Donahue et al., 2012; Kurtén et al., 2016). Therefore, a free parameter b_{add} is included to adjust the effect of oxygen atoms in the molecule b_O and to account for the different functionalities. To obtain this parameter, measured monomer and dimer products with known chemical composition are fitted separately with the group-contribution method SIMPOL (Stolzenburg et al., 2018). A fit to the data yields $b_{add}^{mono} = 0.904$ for HOM monomers and $b_{add}^{di} = 1.139$ for HOM dimers.

355 Consequently, the saturation vapor pressure of any oxidation product measured in the CLOUD chamber can be estimated based on its elemental composition.

In addition, the gas-phase saturation ratio, S_i^*, for each oxidation product can be determined based on the quantitative vapor-phase measurement of the oxidized molecule concentration, $[OVOC_i]$, the molecular mass, m_i, and the associated saturation concentration, C_i^* (Donahue et al., 2013):

$$S_i^* = \frac{[OVOC_i] \cdot \frac{m_i}{N_A}}{C_i^*}. \tag{5}$$

It should be noted that we can only estimate the volatility from the elemental composition, while two molecules with an identical detected mass may have different volatilities depending on their exact chemical structures and functional groups. To account for the dependence of the volatility on temperature, T, the saturation concentration, C_i^0, can be described according to the Clausius-Clapeyron equation:

$$\log_{10} C_i^* (T) = \log_{10} C_i^0 (300K) + \frac{\Delta H_i^{vap}}{R \cdot \ln(10)} \cdot \left(\frac{1}{300K} - \frac{1}{T} \right). \tag{6}$$

According to Donahue et al. (2011) and Epstein et al. (2009), we can approximate the evaporation enthalpy ΔH_i^{vap} as:

$$\Delta H_i^{vap} [kJ\ mol^{-1}] = -5.7 \cdot \log_{10} C_i^* (300K) + 129. \tag{7}$$

Thus, a change in temperature of 15 to 20 K will result in a shift of the volatility bin by one order of magnitude. This study focusses on the oxidation products classified as ELVOC and ULVOC, which will initiate cluster growth and form new particles. However, ELVOC will condense on any particle of any size with negligible re-evaporation, but may not contribute significantly to nucleation itself, while ULVOC in contrast may efficiently nucleate. To account for our incomplete knowledge of the exact chemical structures and functional groups of the oxidation products we assume an overall uncertainty of ±1 bin in the volatility distribution (corresponding to one order of magnitude in $C^*(300K)$).

2.8 HOM formation and its dependence on temperature

Two parameters, AP_{oxrate}^T and γ_{HOM}^T, are used to describe and characterize the overall HOM formation. To account for the different oxidant concentration $[O_3]$ and $[OH]$ among the experiments and the temperature dependence of the initial reaction rate coefficient of α-pinene by these oxidants (Fig. S2(a)) the α-pinene oxidation rate is used as follows:

$$AP_{oxrate}^T [molecule \cdot cm^{-3} \cdot s^{-1}] = k_{(AP+O_3)}^T \cdot [AP] \cdot [O_3] + k_{(AP+OH)}^T \cdot [AP] \cdot [OH]. \tag{8}$$

Here, $[AP]$ and $[O_3]$ are the measured gas-phase concentrations of α-pinene by the PTR3 instrument and ozone by a trace gas monitor, respectively. The IUPAC recommended rate coefficients of the α-pinene ozonolysis reaction ($k_{(AP+O_3)}^T = 8.05 \cdot$

$10^{-16} \cdot e^{-640\text{K}/T} \text{cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$) and the reaction of α-pinene with OH ($k^{\text{T}}_{(\text{AP}+\text{OH})} = 1.2 \cdot 10^{-11} \cdot e^{440\text{K}/T} \text{cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$) are used. The temperature dependence of these rate coefficients is shown in Fig. S2(a) for typical oxidant concentrations used in our experiment. The main sources of OH radicals are the ozonolysis of α-pinene and by UV photolysis of ozone. In dark conditions (UV off), the temperature-dependent ozonolysis rate is a major source of OH radicals with a yield of 80 % (Chew and Atkinson, 1996), with a resulting steady-state OH concentration of $0.5 - 1.6 \cdot 10^6$ molecule cm^{-3}. The formation of OH radicals depends mainly on the absolute humidity in the chamber since singlet D-oxygen, which is formed during the ozone photolysis, is subsequently recombined with H_2O. The OH radical concentration by UV was estimated from dedicated actinometry experiments, forming sulfuric acid, during the same campaign. The OH production by UV yields $1 - 3 \cdot 10^6$ molecule cm^{-3} at $+ 25\,°\text{C}$, while at low temperatures the OH production is comparatively small ($\leq 1 \cdot 10^5$ molecule cm^{-3} at $- 50\,°\text{C}$), due to the lower humidity in the chamber.

The total oxygenated organic fraction [OxOrg] can be estimated as follows

$$\frac{\text{d}\,[OxOrg]}{\text{d}\,t} = AP^{\text{T}}_{\text{oxrate}} - \left(k^{\text{T}}_{\text{dil}} + k^{\text{T}}_{\text{wall}} + k_{\text{CS}}\right) \cdot [OxOrg]. \tag{9}$$

The dilution loss rate k^{T}_{dil} is determined by dividing the total flow into the chamber by its volume ($k^{278\text{K}}_{\text{dil}} \sim 2 \cdot 10^{-4} \text{ s}^{-1}$) equaling total outflow at constant chamber pressure. Since the focus of this study is on compounds that are relevant for nucleation and early growth, we assume that oxygenated organics are irreversibly lost due to condensation on a wall or particles. The chamber wall loss rate was determined to be $k^{278\text{K}}_{\text{wall}} = 2 \cdot 10^{-3} \text{ s}^{-1}$, which is the major loss. An additional loss is due to the condensation sink ($k_{\text{CS}} \sim 0.001$ to $0.1 \cdot 10^{-3} \text{ s}^{-1}$) to particles and the dilution loss ($k^{278\text{K}}_{\text{dil}} \sim 0.2 \cdot 10^{-3} \text{ s}^{-1}$). The total loss rates for oxygenated organics is then $k^{278\text{K}}_{\text{loss}} \sim 2.2 \cdot 10^{-3} \text{ s}^{-1}$. Note that the condensation sink in the CLOUD chamber is lower than in other chamber experiments, where similar experiments have been conducted. Based on the production terms, the cumulative sinks and the total measured [HOM] by the nitrate CI-APi-TOF, the HOM yield ($\gamma^{\text{T}}_{\text{HOM}}$) can be expressed as

$$\gamma^{\text{T}}_{\text{HOM}}[\%] = \left(\frac{\dfrac{\text{d}\,[OxOrg]}{\text{d}\,t} + [\text{HOM}] \cdot \left(k^{\text{T}}_{\text{loss}}\right)}{AP^{\text{T}}_{\text{oxrate}}}\right) \cdot 100 . \tag{10}$$

3 Results and Discussion

3.1 Evolution of gases and particles during an experimental CLOUD run

A typical CLOUD experiment ("run") is performed after establishing a constant level of ozone. Starting from ion-free conditions, α-pinene is added to the chamber at a constant rate, as shown in Fig. 1 at 12:30, 20th October 2018. Due to chemical reactivity, the ozone concentration varied between 34 and 40 ppb$_\text{v}$. As soon as α-pinene was added to the chamber, peroxy radicals (RO_2) and HOMs started to form. In contrast to previous CLOUD campaigns, in CLOUD12 and CLOUD13 the fan

was switched to 100% speed during the addition of α-pinene. The high fan speed increases turbulent mixing in the chamber, and lead to a faster deposition of oxidation products and particles onto the wall (k_{wall}). Consequently, the steady-state

405 concentration of condensable material (ELVOC and ULVOC) were shifted well below the nucleation threshold by increasing the fan speed from its standard value (12%) to 100%. The concentration of the peroxy radicals measured by the CI-APi-TOF, however, is not much affected by the strong fan mixing. Reaction rate constants for highly functionalized RO_2 from α-pinene self and cross reactions are in a range of 1 to $10 \cdot 10^{-11} cm^3$ molecule^{-1} s^{-1} at 300 K (Berndt et al., 2018a). Due to their high reactivity, the lifetime of the RO_2 radicals is mainly determined by chemical loss rates, and relative weakly by the wall loss

410 rate.

After the precursors reached a steady-state concentration (13:23 UTC in Fig. 1), the mixing fans were switched from 100 % to 12 % speed, reducing HOM and cluster wall loss rates by a factor of two to three. Consequently, a new steady-state concentration of α-pinene oxidation product monomers (C_{10}) and dimers (C_{20}) was established on the wall-loss timescale. Due to the increased gas-phase concentration of condensable material, a new-particle formation event was initiated. Molecular

415 clusters started to form and grew into aerosol particles. After the particle formation rate had reached steady state under neutral conditions (J_n), the HVCF inside the chamber was turned off (15:38 – 17:40 UTC in Fig. 1). Due to natural ionization at intensities of ground-level galactic cosmic rays, the ion concentration increased to > 1000 cm^{-3}. Maintaining all other chamber parameters constant, we observed an enhancement of up to two orders of magnitude or more in the nucleation rate of new particles due to ion-induced cluster stabilization (J_{gcr}) (Kirkby et al., 2016).

420 During some stages, the UV light was also turned on to study its effect on the oxidation chemistry by comparing the results with (06:00 – 08:20 UTC in Fig. 1) and without (03:09 – 05:41 UTC in Fig. 1) photochemical reactions under otherwise identical conditions. The particle formation sequence was then repeated at various concentrations of α-pinene and different temperatures over the range of atmospheric interest. In the data analysis, we assume that the particles observed at 1.7 nm mobility diameter are stable against evaporation and serve as a valid proxy for new-particle formation in the chamber.

425 **3.2 Effect of temperature on α-pinene oxidation and HOM formation**

Temperature has a strong effect on peroxy radical isomerization and, consequently, on the production rate of closed-shell oxygenated products. HOM formation is, in principle, controlled by the production rate and lifetime of the precursor peroxy radicals, while the lifetime of the radicals is determined by the competing reaction of the unimolecular autoxidation and the bimolecular terminations. The unimolecular H-shift reaction has a much higher pre-exponential term for the rate constant given

430 by the molecular vibration frequencies compared with that for the bimolecular process, which mainly depends on the bimolecular collision frequency (Praske et al., 2018). However, the higher activation energy barrier of the H-shift reaction partly or fully compensated this. Quantum chemical calculations for different RO_2 radicals from α-pinene oxidation suggest activation energies between 92 and 121 kJ mol^{-1} for the autoxidation process (Rissanen et al., 2015). Because of this high activation energy barrier, temperature has a substantial effect on the intramolecular H-shift and will strongly reduce the

435 autoxidation at lower temperature. In contrast, the temperature dependence of the bimolecular reaction (like molecular dimer

formation) is much weaker or, in some cases, even exhibits a slightly negative dependence. Consequently, the competition at lower temperatures between the termination reaction and the slower unimolecular autoxidation rate influences the oxidation state of the products and their distribution. This temperature-dependence of the α-pinene oxidation was previously observed in the composition of HOM ions charged by cosmic rays in the CLOUD chamber (Frege et al., 2018), and is confirmed here for neutral HOMs and their gas-phase clusters, as shown in Fig. 2. A strong decrease in the mean O/C of the detected oxidation products can be observed as temperature decreases. This decrease in O/C with decreasing temperature was also observed in the particles phase by Kristensen et al. (2017) and Ye et al. (2019) and raises to the question to what extent does the reduction of oxidation also affects new-particle formation?

In general, the HOM production rate is approximately linearly dependent on its initial precursor concentration, while the bimolecular termination of the RO_2 radicals increases with a squared dependency. Molteni et al. (2019) showed that an increase in the α-pinene oxidation rate, and thus an increase in the RO_2 radical concentration, leads to a slightly lower mean O/C of the detected products. However, this effect is rather small at low α-pinene oxidation rates ($< 1 \cdot 10^7$ molecule cm^{-3} s^{-1}) and could not be clearly recognized in our data.

A similar effect of RO_2 radical concentration on the bimolecular terminations can be seen in the mean HOM dimer (C_{20}) to monomer (C_{10}) ratio in Fig. 3. An elevated oxidation rate (AP_{oxrate}^T, color scale) leads to a slightly increased HOM dimer formation. In addition, Fig. 3 shows that the ratio of HOM dimers to monomers is almost unaffected at lower temperatures and depends mostly on the formation of RO_2 radicals. This indicates that the bimolecular termination has only a minor temperature dependence.

At temperatures below 0 °C, the ratio of dimers to monomers is about 20 %. It should be noted that, especially at the lowest temperatures (-50 °C), dimer clusters from traditional oxidation products are formed, e.g., pinonic acid dimers, which do not result from the autoxidation reaction (non-HOM dimer clusters). Due to lower fragmentation in the instrument, these dimers can also be detected by the mass spectrometer. Since we cannot differentiate between them, these dimer clusters are included in the overall gas-phase HOM dimer concentration. However, due to their greatly reduced saturation vapor pressure these dimers also participate in new-particle formation.

At temperatures above 0 °C, the gas-phase HOM monomer concentrations increase more than the HOM dimer concentration, as seen from the decreasing dimer ratio of 20 % to less than 10 % (Fig. 3). This shift can be explained by two factors; volatility and formation rate. At higher temperatures, the saturation vapor pressures of the HOM monomers are sufficient for them to re-evaporate from surfaces (or particles), while the dimers remain essentially irreversibly. This shift in saturation vapor pressure is also consistent with volatility predictions by the VBS model, as shown by the colour code in Fig. 5. This supports our conclusion that the dimer molecules of accretion reactions play a decisive role in new-particle formation, especially at higher temperatures (Mohr et al., 2017; Lehtipalo et al., 2018).

At higher temperatures, the concentration of HO_2 radicals can also strongly increase, as shown by the HO_2 reaction rate coefficients in Fig. S2(b). In addition to their formation by alkoxy radicals (RO), the HO_2 radicals are mainly formed by the reaction of OH radicals with ozone, while the OH radical concentration depends mainly on the ozonolysis reaction rate of α-

470 pinene. Both reactions have a positive temperature dependence. The higher concentration of HO_2 leads to an increased competition with the RO_2-RO_2 self-reaction, which reduces the formation of HOM dimers and increase HOM monomers.

To summarize, the reduced α-pinene ozonolysis rate at lower temperatures leads to lower concentrations of highly-oxygenated organic molecules. Furthermore, due to the strong temperature dependence of the initial H-shift in peroxy radical isomerization, the HOM yield per ozonolysis reaction will also be reduced. As shown in Fig. 4(a), the overall HOM yield

475 drops dramatically as the temperature decreases. From the slopes of the linear fits of HOM measurement by the nitrate CI-APi-TOF (Fig. 4(b)), the total HOM formation of the α-pinene oxidation reactions (O_3 and OH) provides the following HOM yields: $\gamma_{HOM}^{298K} = 6.2$ % at 25 °C, $\gamma_{HOM}^{278K} = 4.7$ % at 5 °C, $\gamma_{HOM}^{263K} = 2.9$ % at -10 °C, $\gamma_{HOM}^{248K} = 1.3$ % at -25 °C and $\gamma_{HOM}^{223K} = 0.7$ % at -50 °C. Whereas these yields at 25 °C and 5 °C are consistent with earlier publications (Ehn et al., 2014; Jokinen et al., 2015; Kirkby et al., 2016), here we extend the HOM yields to much colder temperatures. However, Quéléver et al. (2019)

480 reports a much stronger temperature dependence of the HOM yields with decreasing temperature, which may result from a combination of differences in the experimental and instrumental conditions between the two studies. It is important to note that the HOMs require at least 6 oxygen atoms (Fig. S1); otherwise, the nitrate ionization scheme cannot fully detect them. The unimolecular reaction of the RO_2 radicals proceed at much slower rates at low temperatures, due to the height of the reaction barrier, while the bimolecular termination of the radicals is much less affected by the temperature (Ziemann and

485 Atkinson, 2012; Frege et al., 2018). Therefore, at low temperatures, there is a higher chance that an RO_2 radical will be lost before it reaches a high oxygenation and forms a HOM that can be detected by the nitrate CI-APi-TOF. Besides the termination reactions, relevant loss mechanism of RO_2 radicals are wall loss and uptake on particles. Especially under cold conditions, it becomes more likely that organic molecules with less than 6 oxygen atoms can condense on pre-existing or newly formed particles due to their reduced volatility and can thereby increase the condensation sink compared to high temperatures. As a

490 result, the HOM yield drops significantly depending on the condensation sink, which tends to result in lower determined yields, especially at low temperatures.

3.3 Change in the volatility distribution of α–pinene oxidation products as a function of temperature

Despite the decrease in the total HOM yield, Fig. 4(c) shows an increase in ultra-low volatility products, indicating a more

495 important reduction in the saturation vapor pressure of all oxidation products compared with the reduction of their oxidation states. The definition of HOM is based on chemical makeup, and is temperature-independent, whereas the physical properties, in particular the volatility, do depend on temperature. Therefore, the classification of a molecule as HOM is inadequate for describing its nucleation properties over a wide temperature range. Figure 5 shows a mass defect plot of the oxidation products measured with the nitrate CI-APi-TOF, where the colour code classifies them into the different volatility bands according to

500 the calculated saturation vapor pressures at each temperature. According to Section 2.7., the volatility classes are based on a semi-empirical approach of the 2-dimensional volatility basis set model (2D-VBS model, Donahue et al. (2011)). The area of the markers reflects the measured concentrations. Only experiments with a similar AP_{oxrate}^{T} ($2 - 3 \cdot 10^6$ cm^{-3}) are included,

to ensure comparability in terms of α-pinene oxidation rates and their temperature dependence. The reduction in the mean oxidation state (OS_C) with decreasing temperature is indicated by the shift of the detected products to the left-hand edge (lower masses) of the different homologous sequences of oxidation products, shown by the lines in Fig. 5 and the coloring in Fig. S3. As seen in the 25 °C data in Fig. 5, most HOM monomers belong to the low-volatility (LVOC) or even semi-volatile class (SVOC), which corresponds to the model results of Schervish and Donahue (2019). Therefore, the loss rate of these oxidation products from the gas-phase is low compared to the extremely-low volatility products (ELVOC), which has already been discussed based on the dimer to monomer ratio in Fig. 3. At 25 °C, only the most oxygenated monomers have an extremely low saturation vapor pressure, while none of them meets the ULVOC class. However, the majority of ELVOCs and ULVOCs at 25 °C are HOM dimers. As temperature decreases, the saturation vapor pressures of the HOM monomers shift towards the ELVOC class, despite the decrease in their O/C. Consequently, at -50 °C, almost all the α-pinene oxidation products detected by a nitrate chemical ionisation mass spectrometer are ELVOCs or ULVOCs.

When the concentration of an oxidation product exceeds its saturated vapor concentration, it becomes supersaturated (saturation ratio, $S_i > 1$). The saturation ratio of an oxidation product therefore indicates to what extent this individual product can condense from the gas-phase and contribute to new-particle formation and growth. Based on the measured concentrations and the volatilities derived from the VBS model, the steady-state saturation ratios for each individual compound (a mixture of isomers) can be determined according to Eq. (5). Accordingly, the saturation ratio, S_i, for each identified oxygenated molecule (HOM_i) is given by the area of the marker in Fig. 6. This presentation gives an indication of the ability to form embryonic molecular clusters by the different HOM_i as a function of temperature. Nucleation requires vapours with $S_i \gg 1$, since they must overcome the Kelvin (curvature) barrier. The colour code in Fig. 6 indicates the concentration determined by the nitrate CI-APi-TOF.

Due to their low volatility, the highest oxidation states of the reaction reach high levels of supersaturations. However, their total concentrations are low, so that their abundance is likely to be insignificant for the formation of clusters and subsequent growth. Consequently, the nucleation and initial growth rate may be limited for each individual HOM. Therefore, further support from other condensable gases, such as sulfuric acid and ammonia may be required to grow these clusters to larger sizes before they are lost to walls or pre-existing particles. However, by summing up the saturation ratios, S_i, of all oxidation products, the nucleation ability can be estimated. Based on the -25 °C data in Fig. 6, the proportion of the supersaturated monomers is significantly higher than the supersaturated monomers at higher temperatures. Thus, despite the decrease in oxidation rate, it is expected that the decrease in volatility with decreasing temperature will increase nucleation and initial growth rates at given precursor (see Sect. 3.5).

3.4 CI-APi-TOF HOM detection and the combination to the PTR3 data set

Nitrate chemical ionization mass spectrometry quantitatively measures highly-oxygenated organic molecules with a high O/C (Hyttinen et al., 2017; Ehn et al., 2017). We generally assume that the charging efficiency in the instrument for HOMs is equal

535   to that of sulfuric acid. We also assume that the charging efficiency does not change significantly at different temperatures and humidity (Viggiano et al., 1997), but is mainly influenced by the oxidised volatile organic compounds (OVOCs).

Kinetic simulations by Hyttinen et al. (2017) show that, for nitrate chemical ionization, HOMs are typically charged with the monomer of the reagent ions $HOM_i \cdot NO_3^-$, since this cluster has a stronger binding energy than the dimer or trimer of the reagent ions, $(HNO_3)_{1-2}NO_3^-$. However, we find that a small fraction (~10 %) of the HOMs do cluster with the dimer of the

540   reagent ions. At low temperatures, this fraction increases due to the increased stability of the charged clusters, resulting in an enhanced survival probability of the charged nitrate dimer in the CI-APi-TOF. At -50 °C, some of the detected HOMs even form clusters with the trimer of the reagent ions. The low oxidation states of the α-pinene oxidation products were most efficiently detected with a nitrate dimer ion. Hyttinen et al. (2015) showed that the HOMs formed from cyclohexene oxidation require at least two hydroperoxy groups, or other H-bond-donating groups, in order to form an energetically more favourable

545   cluster than the nitrate dimer or trimer ion cluster. The reason why bonding of low oxidation products to the reagent dimer or trimer ion cluster is preferred is probably the distance between the interacting functional groups of the oxidation product and the size of the reagent ion itself. As shown by Rissanen et al. (2015), the primary ozone attack in α-pinene oxidation results in a ring-opening reaction, leading to a subsequent carbon-centered alkyl-type radical bearing functional groups at both ends of the molecule. Due to the distance between these functional groups, a single nitrate anion can only interact with one of the

550   hydroperoxy groups, while the dimer or trimer reagent ion can bridge this distance and form an energetically more stable ion cluster.

Since both the oxidation state and the saturation ratio of the oxidation products depend on temperature, some of the less oxygenated but nevertheless highly supersaturated OVOCs may be underestimated or even undetected by a nitrate CI-APi-TOF. Figure 6 should therefore be considered as a lower limit at low temperatures (-25 °C and -50 °C), where most oxidation

555   products have a low O/C. Furthermore, the saturation ratio of the oxygenated dimer products (C_{20}) may be underestimated, even at relatively high temperatures, due to their generally lower O/C.

The PTR3 instrument, which uses an H_3O^+ reagent ions, is well-suited to measure the less highly-oxygenated products. Figure S1 shows a comparison of the homologous groups of $C_{10}H_{14,16}O_x$ measured with both instruments. The concentration of dimer molecules did not exceed the limit of detection (LOD, $8 \cdot 10^5$ cm^{-3}) of the PTR3 instrument. Therefore, no direct statement

560   can be made here about the concentration and saturation ratio of the less oxygenated dimer products. However, as the concentrations are below the LOD, even though they can initiate new-particle formation, they may provide only a small amount of condensable mass to promote early growth. As pointed out in Stolzenburg et al. (2018), the concentration of products with an O/C of 0.6 to 0.7 is in good agreement between the two instruments, as shown in Fig. S1, while the nitrate CI-APi-TOF lacks the less-oxygenated products. The different colour segments on the lower axes, however, show that the nitrate CI-APi-

565   TOF can quantitatively measure especially the ELVOC and ULVOC products at mixing ratios in the ppq$_v$ range. Since the ULVOCs are the most effective molecules for nucleation (Schervish and Donahue, 2019), the nitrate ionization instrument provides a complete molecular understanding of new-particle formation over a wide range of tropospheric temperatures.

Figure 7 shows all oxidation products distributed according to their volatility (in bins of factors of ten) measured by the nitrate CI-APi-TOF and PTR3 instruments at each temperature. While these data are consistent with the results of Stolzenburg et al. (2018), the present study covers a wider temperature range from -50 °C to 25 °C. The dashed lines in Fig. 7 represent the cumulative concentrations, starting with the lowest volatility products measured in the nitrate CI-APi-TOF. The solid lines include the less highly-oxygenated products measured in the PTR3 instrument, assuming the same volatility products for each bin. While the nitrate CI-APi-TOF can completely quantify the ULVOC and ELVOC class at temperatures ≥ -25 °C, the molecules measured by the PTR3 contribute to the total ELVOC class at the lowest temperature (-50 °C). The ULVOC class is fully represented by the nitrate CI-APi-TOF and is anti-correlated with temperature, as already shown in Fig. 4(c). Furthermore, Fig. 6 shows that, at low temperature (-50 °C and -25 °C), almost all oxidation products detected by the nitrate CI-APi-TOF are supersaturated at the indicated precursor concentration and can form stable embryonic clusters and small particles. Due to the large decrease in volatility at these low temperatures, a large fraction of the less oxygenated molecules also falls in the ELVOC class. While the ULVOC and ELVOC molecules are evenly distributed over a wide range of volatility bins, an uncertainty in the saturation concentration of one-order of magnitude would not alter this conclusion.

3.5 Effect of temperature on pure biogenic nucleation

The decrease in HOM volatility at reduced temperatures has a strong effect on the particle formation rates at 1.7 nm ($J_{1.7}$), as seen in Fig. 8. Despite the decrease in autoxidation rate of the peroxy radicals and the associated reduction in HOM yields at lower temperatures, the reduced volatility counteracts this effect and leads to an overall increase in particle formation rate. Figure 8(a) presents the particle formation rate as a function of the total HOM concentration measured by the nitrate CI-APi-TOF. This figure shows that the same or higher nucleation rates are reached for HOM concentrations a factor 100 lower at -50 °C compared to 25 °C. At low temperatures, molecules with less oxygenation can drive the same formation rate as more higly oxygenated molecules at higher temperatures. However, the chemical definition of HOM does not reflect the saturation vapor pressure of the oxidation products. This representation therefore underestimates the condensable mass at low temperatures where the high gas-phase concentration of the less-oxygenated products can contribute to sufficient supersaturation and new-particle formation (see Sect. 3.4).

The total HOM concentration depends mainly on the C_{10} monomer concentration, while the C_{20} dimers only account for 10 % to 20 % of that total, depending on the experimental conditions (e.g., HO_2:RO_2, temperature, precursor concentration). Whereas most of the C_{10} monomers have a saturation vapor pressure in the range of SVOC to LVOC at high temperatures (> 5 °C), at low temperature (-50 °C) almost all of them belong to the ELVOC and even ULVOC class. At higher temperatures, dimerization (C_{20} formation) can be the rate-limiting step for new-particle formation since the volatility not only depends on the polarity of the molecule (functional groups), but also on its size (m/Z). Thus, a major fraction of the HOM dimers fall in the ELVOC to ULVOC range, even at 25 °C (see also Fig. 5). Therefore, the ULVOC class may provide a more appropriate proxy to parameterize nucleation rates from monoterpene oxidation. Note that dimer clusters (non-covalently bonded clusters)

600   may form at lower temperatures (< -25 °C), which, despite their low degree of oxidation, also have a low saturation vapor pressure, and so may contribute both to new-particle formation and to the total dimer concentration.

The continuous decrease in the autoxidation rate coefficient with decreasing temperatures reduces the HOM concentrations in Fig. 8(a), for the same precursor concentrations used at higher temperatures. Therefore, Fig. 8(b), shows the new-particle formation rate as a function of the sum of all molecules in the ULVOC range. The evaporation rate of a cluster containing two

605   ULVOCs will conceptually be related to the average volatility of those two molecules. Due to the sufficiently low saturation vapor pressure, we consider the ULVOC range to be the volatility range where it is increasingly likely that a cluster will grow, not evaporate, and thus nucleate. This is consistent with the result shown in Fig. 8(b) as it unifies the new-particle formation rates at all temperatures and shows that ULVOC concentrations provides a better proxy compared to the total HOM concentration. Ideally, all lines of the nucleation rates for the different temperatures would coincide into a single line in Fig.

610   8b. However, the nucleation rate will still likely be a function of volatility within the ULVOC (and partially the ELVOC) range. Based on the previous discussion, however, we can infer that both mass spectrometers underestimate some low-volatility products, especially at the two lowest temperatures. Fragmentation and/or clustering of the oxygenated molecules inside the mass spectrometers could also lead to misinterpretation of their volatility (Passananti et al., 2019). In addition, the temperature dependence of the volatility classification is subject to larger uncertainty as the temperature decreases. Thus, we estimate that

615   the bin assignment is uncertain by about one bin (a factor of 10 in volatility). Further, the stabilizing effect of ions is not considered in the VBS model. Under ion-free condition, the pure biogenic nucleating clusters are generally less stable against evaporation. It is possible that the stabilizing effect of an ion essentially shifts the relevant volatility range by some number of bins and results in a larger uncertainty at warm temperatures, while this effect is less apparent at colder temperatures in absence of any ions for the same ULVOC concentration. The probability that a cluster grows under ion-free conditions, due to an

620   arriving molecule, is strongly dependent on its volatility compared to GCR conditions. Therefore, the exact volatility distribution can also be important within the ULVOC class, especially for the neutral conditions.

Figure 8(c) shows the neutral particle formation rates at 1.7 nm (J_n) and Fig. 8(d) the galactic cosmic ray particle formation rates ($J_{gcr} = J_n + J_{iin}$) as a function of the α-pinene oxidation rate (AP^T_{oxrate}). This representation accounts for the change in the oxidation rate coefficient of α-pinene with respect to temperature and is independent of the concentration of the oxidation

625   products and their uncertainties in detection by the instrument as well as the HOM definition and the volatility classification. Moreover, the representation of the nucleation rates as a function of the total concentrations, not volume mixing ratios, removes the influence of the pressure effects on the molecule/cluster collision rates at different atmospheric altitudes.

Higher precursor concentrations lead to an increase in supersaturated mass of the mixture, and thus, to an increased new-particle formation rate. This can be seen from the positive slope of the power-law fits (lines) from the neutral new-particle

630   formation rates (J_n) in Fig. 8(c). The slopes are sensitive to the interplay between particle growth rates and wall loss rates, as well as to the molecular composition of the critical clusters (Ehrhart and Curtius, 2013). In addition to the poorly quantified effect of temperature on the operation of condensation particle counters, including the PSM (Wimmer et al., 2015), another point to consider is that the new-particle formation rates under ion-free conditions, especially at warmer temperatures, are low.

Due to very low particle concentrations in the chamber, the data are close to the lower detection limit of the measurable

635   formation rates. Both factors lead to a higher uncertainty in the nucleation rate calculation (Dada et al., 2020).

While there is no significant change in the new-particle formation rate from +25 °C to -10 °C, a sharp increase of up to two orders of magnitude in J_n occurs at lower temperatures. This non-linear behavior of the nucleation rates with decreasing temperature can be explained by the increased stability of the neutral clusters against evaporation without any stabilizing agent (*e.g.,* ions), along with the strong temperature shift of the volatility classes. At low temperatures, more products fall in the

640   ULVOC range and the increased concentration of LVOCs may enhance nucleation and certainly growth (Fig. 8). Due to the strong reduction in volatility, also first-order oxidation products (OVOCs) will take part in the initial growth of nucleating clusters at much smaller cluster sizes, while their oxidation rate is much less affected by the temperature compared to the autoxidation rate.

645   HOMs have high electrical polarizability, so that ions can stabilize the embryonic clusters and also lead to a higher collision rate with other HOMs. Figure 8(d) shows that ions (GCR conditions) enhance the new-particle formation rates by up to two orders of magnitude or more compared to neutral conditions, depending on the precursor concentrations and temperature. At higher levels of AP^T_{oxrate}, the neutral and the ion-induced cluster formation rates eventually converge (horizontal grey line), as the ion-induced nucleation rate is limited by the total ion-production rate (~4.1 ion pairs cm^{-3} s^{-1}) resulting from the galactic

650   cosmic ray intensities at ground level. This comparison shows that the neutral HOM clusters are relatively unstable, and, especially at high temperatures, have significant evaporation rates, which demonstrates the strong sensitivity of the pure biogenic nucleation to environmental conditions and the concentration of ions (Kirkby et al., 2016). The oxidation rate (AP^T_{oxrate}) at which J_n and J_{gcr} converge, however, depends strongly on temperature. At -50 °C the neutral and GCR nucleation rates almost coincide over the range of experimental measurements, reflecting the increased stability of the neutral clusters at

655   low temperatures. This strong temperature sensitivity leads to a high nucleation potential of biogenic molecular clusters especially in cold regions (i.e., convective updraft of monoterpenes to higher altitudes). While our experiments were carried out at ground-level pressure (~970 hPa), we expect that our concentration-based new-particle formation rates also apply to the upper troposphere within their uncertainties. The size and number of heavy atoms in the nucleating clusters containing two ULVOCs allow for the efficient distribution of the cluster formation energy in the form of vibrational and rotational excitation

660   within the cluster. Therefore, we do not anticipate a strong pressure dependence throughout the troposphere for biogenic nucleation involving HOMs.

However, there might be other factors affecting the new-particle formation in the upper troposphere. These include the increased ionization capacity, the absolute humidity, the coagulation sink and the availability of condensable gases for particle growth. Also, some chemical reactions are pressure dependent, but compared to their temperature dependence this is mostly a

665   minor effect. Moreover, we believe that most of the HOMs are likely produced in the vicinity of their precursor sources, due to their high reactivity with oxidants, and then transported to higher altitudes by convective updrafts (Williamson et al., 2019) or are produced by second-generation chemistry. However, this is speculative and needs to be investigated in future studies.

3.6 Influence of UV light on α-pinene oxidation and new-particle formation

In addition to ozone, OH radicals oxidise α-pinene and form HOM, albeit with a lower yield. The lower yield results from the inability of OH to break the 6-carbon ring of the α-pinene molecule, resulting in steric hindrance of the subsequent RO_2 isomerization (Ehn et al., 2014; Berndt et al., 2016; Kirkby et al., 2016). This results in a higher energy barrier for formation of the intermediate alkyl radicals ($RO^{\bullet}$) and thus slows down the H-shift rate, which decrease highly oxidised products with low volatilities.

Figure S4 presents a mass defect plot showing the change between the UV-on and UV-off conditions of the last two GCR stages of the run shown in Fig. 1. The elevated OH radical concentration induced by the UV light leads to products with higher H/C ratio than found in dark ozonolysis experiments (Fig. S5). The initial reaction of OH radicals with α-pinene forms reactive RO_2 intermediate radicals of the chemical notation $C_{10}H_{17}O_{odd}$, which also undergo several autoxidation steps (Berndt et al., 2018a). The bimolecular termination of this propagation chain leads to more HOM monomer molecules of the chemical formula $C_{10}H_{16,18}O_x$ and subsequently more HOM dimer molecules of the type $C_{20}H_{32,34}O_x$ (red and green trace, Fig. S6(d)-(e)), while products of the type $C_{10}H_{14}O_x$ and $C_{20}H_{30}O_x$ are reduced (blue trace, Fig. S6(d)-(e)). We note that nitrate ionization may have a reduced charging efficiency for HOMs formed from OH oxidation compared with those from ozonolysis, suggesting a possible measurement bias (Berndt et al., 2016).

Increasing the OH concentration in α-pinene oxidation experiments not only alters the molecular structure of the initial RO_2 radicals but also influences the lifetime of the RO_2 radicals and the HOM yield. The increase in OH affects the HO_X cycle and may increase HO_2 radicals. The HO_2 radicals will react with intermediate RO_2 radicals and terminate the autoxidation at an earlier stage, leading mainly to closed-shell products of the type $C_{10}H_{16}O_{2+2n}$. The HO_2 radicals also compete with the RO_2-RO_2 cross-reaction and so reduce dimer formation. Another source of HOM monomers may also be the subsequent reaction of primary α-pinene oxidation products (e.g., pinanediol) with OH radicals (Schobesberger et al., 2013; Ye et al., 2018).

As a result, the nucleation rates can be significantly influenced by the HO_x relative to RO_2 concentrations (HO_2/RO_2 ratio), especially at higher temperature where the dimerization may be the rate-limiting step of pure biogenic nucleation. Compared to the atmosphere, in chamber experiments the RO_2/HO_2 ratio is shifted in the direction of the RO_2 radical (Molteni et al., 2019). The difference between dark conditions and UV illumination may therefore be less pronounced in our study than in the real atmosphere. Keeping this caveat in mind, no pronounced variation in the overall HOM yield between UV illumination and its absence can be seen in Fig. 4. Furthermore, no photochemical reaction (degradation) seems to affect the total HOM yield. In summary, Fig. 8 shows similar nucleation rates between dark ozonolysis and UV-illuminated conditions, as indicated by the orange-framed symbols.

4 Conclusions

Measurements made in the CERN CLOUD chamber with two mass spectrometers – one using nitrate chemical ionization and the other a complementary proton transfer reaction – have provided a comprehensive molecular understanding of new-particle

formation from α-pinene between -50 °C and 25 °C. Using the mass spectrometry measurements, we performed two-dimensional volatility basis set calculations and determined the volatility distributions at each temperature. While at high temperatures the HOM dimers appear to be crucial for nucleation, at lower temperatures the HOM monomers and less-oxygenated products also contribute.

Unimolecular autoxidation is key to the formation of HOMs from biogenic precursors. Intramolecular isomerization reactions, including the H-atom shift that leads to increased oxygenation, are found to be very sensitive to temperature. We observe a continuous reduction in the oxidation level (O/C ratio) and yield of HOMs as the temperature decreases from 25 °C to -50 °C. However, despite the lower oxidation level and lower yield of HOMs, our measurements show that the reduction of volatility more than compensates and, consequently, that the nucleation rates increase at lower temperatures.

This study demonstrates that chemical ionization mass spectrometers can measure the elemental composition and distribution of biogenic molecules over a wide range of oxidation states (O/C ratio) and at temperatures as low as -50 °C. PTR and nitrate chemical ionization mass spectrometers are mandatory in future field and aircraft campaigns to study these processes in the atmosphere. Our results underscore the potential importance of biogenic nucleation throughout pristine regions of the atmosphere, and especially in cold upper tropospheric environments exposed to convective outflows.

Data availability. Data related to this article are available upon request to the corresponding authors.

Author contributions. M.Sim., L.D., M.H., W.S., D.S., L.F., A.C.W., B.R., A.K., X.H., J.A., R.B., A.Ba., A.Be., F.B., S.Brä., L.C., D.C., B.C., A.D., J.Du., I.E., H.F., C.F., L.G., H.G., M.G., J.H., V.H., C.K., W.K., H.L., C.P.L., K.L., M.L., H.M., H.E.M., G.M., B.M., U.M., A.On., E.P., T.P., J.P., M.P., L.L.Q., M.P.R., S.Scho., S.Schu., J.S., M.Sip., G.S., Y.S., Y.J.T., A.R.T., M.V., A.V., R.W., M.W., D.S.W., Y.W., S.K.W., Y.W., C.Y., P.Y., Q.Y., M.Z., X.Z., R.C.F., R.V., P.M.W. and J.K. contributed to the development of the CLOUD facility and analysis instruments. M.Sim., L.D., M.H., W.S., D.S., L.F., A.C.W., B.R., X.H., J.A., R.B., A.Ba., P.S.B., L.B., A.Be., F.B., S.Brä., S.Bri., L.C., A.D., D.C.D., J.Du., I.E., H.F., L.G., H.G., M.G., J.H., V.H., C.R.H., C.K., W.K., H.L., C.P.L., K.L., M.L., H.E.M., G.M., R.M., B.M., U.M., L.N., W.N., A.Oj., E.P., J.P., L.L.Q., A.R., M.P.R., S.Scha., S.Schu., J.S., G.S., Y.S., C.T., Y.J.T., A.R.T., M.V., A.V., R.W., D.S.W., Y.W., S.K.W., Y.W., M.X., C.Y., P.Y., M.Z., P.M.W. and J.K. collected the data. M.Sim., L.D., M.H., W.S., L.F., A.C.W., B.R., L.G., C.K., G.M., B.M., S.K.W., P.Y. and R.V. performed modelling and analyzed the data. M.Sim., L.D., M.H., W.S., D.S., L.F., A.C.W., B.R., A.K., H.G., B.M., U.M., T.P., M.P.R., P.Y., U.B., J.Do., R.C.F., A.H., M.K., D.R.W., N.M.D., J.K. and J.C. were involved in the scientific discussion and interpretation of the results. M.Sim., L.D., M.H., W.S., A.K., A.V., U.B., J.Do., R.C.F., N.M.D., J.K. and J.C. contributed to writing and editing of the manuscript.

Competing interests. The author declares that there is no conflict of interest.

Acknowledgements. We thank the European Organization for Nuclear Research (CERN) for supporting CLOUD with important technical and financial resources, and for providing a particle beam from the CERN Proton Synchrotron. We thank P. Carrie, L.-P. De Menezes, J. Dumollard, K. Ivanova, F. Josa, I. Krasin, T. Keber, R. Kristic, A. Laassiri, O. S. Maksumov, B. Marichy, H. Martinati, S. V. Mizin, R. Sitals, A. Wasem and M. Wilhelmsson for their contributions to the experiment.

Financial support. This research has received funding from the German Federal Ministry of Education and Research, CLOUD-12 (01LK1222A) and CLOUD-16 (01LK1601A), the EC Seventh Framework Programme and European Union's Horizon 2020 programme (Marie Skłodowska Curie ITNs no. 316662 "CLOUD-TRAIN" and no. 764991 "CLOUD-MOTION", MSCA-IF no. 656994 "nano-CAVa", MC-COFUND grant no. 600377, European Research Council (ERC; projects no. 692891 "DAMOCLES", no. 638703 "COALA", no. 616075 "NANODYNAMITE", no. 335478 "QAPPA", no. 742206 "ATM-GP", no. 714621 "GASPARCON"), the Swiss National Science Foundation (projects no. 20020_152907, 200020_172602, 20FI20_159851, 20FI20_172622), the Academy of Finland (Center of Excellence no. 307331, projects 299574, 296628, 306853, 304013, 310682), the Finnish Funding Agency for Technology and Innovation, the Väisälä Foundation, the Nessling Foundation, the Austrian Science Fund (FWF; project no. J3951-N36, project no. P27295-N20), the Austrian research funding association (FFG, project no. 846050), the Portuguese Foundation for Science and Technology (project no. CERN/FIS-COM/0014/2017), the Swedish Research Council Formas (project number 2015-749), Vetenskapsrådet (grant 2011-5120), the Presidium of the Russian Academy of Sciences and Russian Foundation for Basic Research (grants 08-02-91006-CERN, 12-02-91522-CERN), the U.S. National Science Foundation (grants AGS1136479, AGS1447056, AGS1439551, CHE1012293, AGS1649147, AGS1602086, AGS1801280, AGS1801329, AGS1801574, AGS1801897), the Wallace Research Foundation, the US Department of Energy (grant DE-SC0014469), the NERC GASSP project NE/J024252/1m, the Royal Society (Wolfson Merit Award), United Kingdom Natural Environment Research Council grant NE/K015966/1, Dreyfus Award EP-11-117, the French National Research Agency the Nord-Pas de Calais, European Funds for Regional Economic Development (Labex-Cappa grant ANR-11-LABX-0005-01).

References

Almeida, J., Schobesberger, S., Kürten, A., Ortega, I. K., Kupiainen-Määttä, O., Praplan, A. P., Adamov, A., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Donahue, N. M., Downard, A., Dunne, E., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Guida, R., Hakala, J., Hansel, A., Heinritzi, M., Henschel, H., Jokinen, T., Junninen, H., Kajos, M., Kangasluoma, J., Keskinen, H., Kupc, A., Kurtén, T., Kvashin, A. N., Laaksonen, A., Lehtipalo, K., Leiminger, M., Leppä, J., Loukonen, V., Makhmutov, V., Mathot, S., McGrath, M. J., Nieminen, T., Olenius, T., Onnela, A., Petäjä, T., Riccobono, F., Riipinen, I., Rissanen, M., Rondo, L., Ruuskanen, T., Santos, F. D., Sarnela, N., Schallhart, S., Schnitzhofer, R., Seinfeld, J. H., Simon, M., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A., Tröstl, J., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y., Virtanen, A., Vrtala, A., Wagner, P. E., Weingartner, E., Wex, H., Williamson, C.,

Wimmer, D., Ye, P., Yli-Juuti, T., Carslaw, K. S., Kulmala, M., Curtius, J., Baltensperger, U., Worsnop, D. R., Vehkamäki, H., and Kirkby, J.: Molecular understanding of sulphuric acid–amine particle nucleation in the atmosphere, Nature, 502, 359, 2013.

770 Ball, S. M., Hanson, D. R., Eisele, F. L., and McMurry, P. H.: Laboratory studies of particle nucleation: Initial results for H2SO4, H2O, and NH3 vapors, Journal of Geophysical Research: Atmospheres, 104, 23709-23718, 1999.

Berndt, T., Mentler, B., Scholz, W., Fischer, L., Herrmann, H., Kulmala, M., and Hansel, A.: Accretion Product Formation from Ozonolysis and OH Radical Reaction of α-Pinene: Mechanistic Insight and the Influence of Isoprene and Ethylene, Environmental science & technology, 2018a. 2018a.

Berndt, T., Richters, S., Jokinen, T., Hyttinen, N., Kurtén, T., Otkjær, R. V., Kjaergaard, H. G., Stratmann, F., Herrmann, H.,
775 Sipilä, M., Kulmala, M., and Ehn, M.: Hydroxyl radical-induced formation of highly oxidized organic compounds, Nature Communications, 7, 13677, 2016.

Berndt, T., Richters, S., Kaethner, R., Voigtländer, J., Stratmann, F., Sipilä, M., Kulmala, M., and Herrmann, H.: Gas-Phase Ozonolysis of Cycloalkenes: Formation of Highly Oxidized RO2 Radicals and Their Reactions with NO, NO2, SO2, and Other RO2 Radicals, The Journal of Physical Chemistry A, 119, 10336-10348, 2015.

780 Berndt, T., Scholz, W., Mentler, B., Fischer, L., Herrmann, H., Kulmala, M., and Hansel, A.: Accretion Product Formation from Self- and Cross-Reactions of RO2 Radicals in the Atmosphere, Angewandte Chemie International Edition, 57, 3820-3824, 2018b.

Bernhammer, A. K., Breitenlechner, M., Keutsch, F. N., and Hansel, A.: Technical note: Conversion of isoprene hydroxy hydroperoxides (ISOPOOHs) on metal environmental simulation chamber walls, Atmos. Chem. Phys., 17, 4053-4062,
785 2017.

Bianchi, F., Garmash, O., He, X., Yan, C., Iyer, S., Rosendahl, I., Xu, Z., Rissanen, M. P., Riva, M., and Taipale, R.: The role of highly oxygenated molecules (HOMs) in determining the composition of ambient ions in the boreal forest, Atmospheric Chemistry and Physics, 17, 13819-13831, 2017.

Bianchi, F., Kurtén, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crounse, J. D., Wennberg, P. O., and
790 Mentel, T. F.: Highly oxygenated organic molecules (HOM) from gas-phase autoxidation involving peroxy radicals: A key contributor to atmospheric aerosol, Chemical reviews, 119, 3472-3509, 2019.

Bianchi, F., Tröstl, J., Junninen, H., Frege, C., Henne, S., Hoyle, C. R., Molteni, U., Herrmann, E., Adamov, A., and Bukowiecki, N.: New particle formation in the free troposphere: A question of chemistry and timing, Science, 352, 1109-1112, 2016.

795 Breitenlechner, M., Fischer, L., Hainer, M., Heinritzi, M., Curtius, J., and Hansel, A.: PTR3: An Instrument for Studying the Lifecycle of Reactive Organic Carbon in the Atmosphere, Analytical Chemistry, 89, 5824-5831, 2017.

Chew, A. A. and Atkinson, R.: OH radical formation yields from the gas-phase reactions of O3 with alkenes and monoterpenes, Journal of Geophysical Research: Atmospheres, 101, 28649-28653, 1996.

Crounse, J. D., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., and Wennberg, P. O.: Autoxidation of Organic Compounds
800 in the Atmosphere, The Journal of Physical Chemistry Letters, 4, 3513-3520, 2013.

Dada, L., Lehtipalo, K., Kontkanen, J., Nieminen, T., Baalbaki, R., Ahonen, L., Duplissy, J., Yan, C., Chu, B., Petäjä, T., Lehtinen, K., Kerminen, V.-M., Kulmala, M., and Kangasluoma, J.: Formation and growth of sub-3-nm aerosol particles in experimental chambers, Nature Protocols, 15, 1013-1040, 2020.

Dias, A., Ehrhart, S., Vogel, A., Mathot, S., Onnela, A., Almeida, J., Kirkby, J., Williamson, C., and Mumford, S.: Temperature
805 uniformity in the CERN CLOUD chamber, Atmos. Meas. Tech., 10, 5075-5088, 2017.

Donahue, N. M., Epstein, S., Pandis, S. N., and Robinson, A. L.: A two-dimensional volatility basis set: 1. organic-aerosol mixing thermodynamics, Atmospheric Chemistry and Physics, 11, 3303-3318, 2011.

Donahue, N. M., Kroll, J., Pandis, S. N., and Robinson, A. L.: A two-dimensional volatility basis set–Part 2: Diagnostics of organic-aerosol evolution, Atmospheric Chemistry and Physics, 12, 615-634, 2012.

810 Donahue, N. M., Ortega, I. K., Chuang, W., Riipinen, I., Riccobono, F., Schobesberger, S., Dommen, J., Baltensperger, U., Kulmala, M., and Worsnop, D. R.: How do organic vapors contribute to new-particle formation?, Faraday discussions, 165, 91-104, 2013.

Duplissy, J., Merikanto, J., Franchin, A., Tsagkogeorgas, G., Kangasluoma, J., Wimmer, D., Vuollekoski, H., Schobesberger, S., Lehtipalo, K., Flagan, R. C., Brus, D., Donahue, N. M., Vehkamäki, H., Almeida, J., Amorim, A., Barmet, P., Bianchi,
815 F., Breitenlechner, M., Dunne, E. M., Guida, R., Henschel, H., Junninen, H., Kirkby, J., Kürten, A., Kupc, A., Määttänen, A., Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Praplan, A. P., Riccobono, F., Rondo, L., Steiner, G., Tome, A., Walther, H., Baltensperger, U., Carslaw, K. S., Dommen, J., Hansel, A., Petäjä, T., Sipilä, M., Stratmann, F., Vrtala, A., Wagner, P. E., Worsnop, D. R., Curtius, J., and Kulmala, M.: Effect of ions on sulfuric acid-water binary particle formation: 2. Experimental data and comparison with QC-normalized classical nucleation theory, Journal of Geophysical
820 Research: Atmospheres, 121, 1752-1775, 2016.

Dusek, U., Frank, G., Hildebrandt, L., Curtius, J., Schneider, J., Walter, S., Chand, D., Drewnick, F., Hings, S., and Jung, D.: Size matters more than chemistry for cloud-nucleating ability of aerosol particles, Science, 312, 1375-1378, 2006.

Ehn, M., Berndt, T., Wildt, J., and Mentel, T.: Highly Oxygenated Molecules from Atmospheric Autoxidation of Hydrocarbons: A Prominent Challenge for Chemical Kinetics Studies, International Journal of Chemical Kinetics, 49, 821-831, 2017.

Ehn, M., Thornton, J. A., Kleist, E., Sipila, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurten, T., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T., Petaja, T., Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol, Nature, 506, 476-479, 2014.

Ehrhart, S. and Curtius, J.: Influence of aerosol lifetime on the interpretation of nucleation experiments with respect to the first nucleation theorem, Atmos. Chem. Phys., 13, 11465-11471, 2013.

Eisele, F. and Tanner, D.: Measurement of the gas phase concentration of H2SO4 and methane sulfonic acid and estimates of H2SO4 production and loss in the atmosphere, Journal of Geophysical Research: Atmospheres, 98, 9001-9010, 1993.

Epstein, S. A., Riipinen, I., and Donahue, N. M.: A semiempirical correlation between enthalpy of vaporization and saturation concentration for organic aerosol, Environmental science & technology, 44, 743-748, 2009.

Frege, C., Ortega, I. K., Rissanen, M. P., Praplan, A. P., Steiner, G., Heinritzi, M., Ahonen, L., Amorim, A., Bernhammer, A. K., Bianchi, F., Brilke, S., Breitenlechner, M., Dada, L., Dias, A., Duplissy, J., Ehrhart, S., El-Haddad, I., Fischer, L., Fuchs, C., Garmash, O., Gonin, M., Hansel, A., Hoyle, C. R., Jokinen, T., Junninen, H., Kirkby, J., Kürten, A., Lehtipalo, K., Leiminger, M., Mauldin, R. L., Molteni, U., Nichman, L., Petäjä, T., Sarnela, N., Schobesberger, S., Simon, M., Sipilä, M., Stolzenburg, D., Tomé, A., Vogel, A. L., Wagner, A. C., Wagner, R., Xiao, M., Yan, C., Ye, P., Curtius, J., Donahue, N. M., Flagan, R. C., Kulmala, M., Worsnop, D. R., Winkler, P. M., Dommen, J., and Baltensperger, U.: Influence of temperature on the molecular composition of ions and charged clusters during pure biogenic nucleation, Atmos. Chem. Phys., 18, 65-79, 2018.

Fu, Y., Xue, M., Cai, R., Kangasluoma, J., and Jiang, J.: Theoretical and experimental analysis of the core sampling method: Reducing diffusional losses in aerosol sampling line, Aerosol Science and Technology, 2019. 1-9, 2019.

Gordon, H., Kirkby, J., Baltensperger, U., Bianchi, F., Breitenlechner, M., Curtius, J., Dias, A., Dommen, J., Donahue, N. M., and Dunne, E. M.: Causes and importance of new particle formation in the present-day and pre-industrial atmospheres, Journal of Geophysical Research: Atmospheres, 2017. 2017.

Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J.

D., Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic aerosol: current and emerging issues, Atmos. Chem. Phys., 9, 5155-5236, 2009.

855   Hanson, D. and Eisele, F.: Diffusion of H2SO4 in humidified nitrogen: Hydrated H2SO4, The Journal of Physical Chemistry A, 104, 1715-1719, 2000.

Heinritzi, M., Simon, M., Steiner, G., Wagner, A. C., Kürten, A., Hansel, A., and Curtius, J.: Characterization of the mass-dependent transmission efficiency of a CIMS, Atmos. Meas. Tech., 9, 1449-1460, 2016.

Hyttinen, N., Kupiainen-Määttä, O., Rissanen, M. P., Muuronen, M., Ehn, M., and Kurtén, T.: Modeling the charging of highly

860   oxidized cyclohexene ozonolysis products using nitrate-based chemical ionization, The Journal of Physical Chemistry A, 119, 6339-6345, 2015.

Hyttinen, N., Rissanen, M. P., and Kurtén, T.: Computational Comparison of Acetate and Nitrate Chemical Ionization of Highly Oxidized Cyclohexene Ozonolysis Intermediates and Products, The Journal of Physical Chemistry A, 121, 2172-2179, 2017.

865   Jimenez, J. L., Canagaratna, M., Donahue, N., Prevot, A., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., and Ng, N.: Evolution of organic aerosols in the atmosphere, Science, 326, 1525-1529, 2009.

Jokinen, T., Berndt, T., Makkonen, R., Kerminen, V.-M., Junninen, H., Paasonen, P., Stratmann, F., Herrmann, H., Guenther, A. B., and Worsnop, D. R.: Production of extremely low volatile organic compounds from biogenic emissions: Measured yields and atmospheric implications, Proceedings of the National Academy of Sciences, 112, 7123-7128, 2015.

870   Jokinen, T., Sipilä, M., Junninen, H., Ehn, M., Lönn, G., Hakala, J., Petäjä, T., Mauldin Iii, R. L., Kulmala, M., and Worsnop, D. R.: Atmospheric sulphuric acid and neutral cluster measurements using CI-APi-TOF, Atmos. Chem. Phys., 12, 4117-4125, 2012.

Kerminen, V.-M., Chen, X., Vakkari, V., Petäjä, T., Kulmala, M., and Bianchi, F.: Atmospheric new particle formation and growth: review of field observations, Environmental Research Letters, 13, 103003, 2018.

875   Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné, S., Ickes, L., Kürten, A., Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S., Tsagkogeorgas, G., Wimmer, D., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S., Hansel, A., Hauser, D., Jud, W., Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R., Makhmutov, V., Mathot, S., Mikkilä, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A., Pereira, P., Petäjä, T., Schnitzhofer, R., Seinfeld, J. H.,

880   Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A., Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E., Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger, U., and Kulmala, M.: Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol nucleation, Nature, 476, 429, 2011.

Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M., Yan, C., Almeida, J., Tröstl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim, A., Bernhammer, A.-K., Bianchi, F.,

885   Breitenlechner, M., Brilke, S., Chen, X., Craven, J., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Kim, J., Krapf, M., Kürten, A., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Molteni, U., Onnela, A., Peräkylä, O., Piel, F., Petäjä, T., Praplan, A. P., Pringle, K., Rap, A., Richards, N. A. D., Riipinen, I., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Scott, C. E., Seinfeld, J. H., Sipilä, M., Steiner, G., Stozhkov, Y., Stratmann, F., Tomé, A., Virtanen, A., Vogel, A. L., Wagner, A. C.,

890   Wagner, P. E., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R., Baltensperger, U., Kulmala, M., Carslaw, K. S., and Curtius, J.: Ion-induced nucleation of pure biogenic particles, Nature, 533, 521-526, 2016.

Knopf, D. A., Pöschl, U., and Shiraiwa, M.: Radial diffusion and penetration of gas molecules and aerosol particles through laminar flow reactors, denuders, and sampling tubes, Analytical chemistry, 87, 3746-3754, 2015.

895   Kristensen, K., Jensen, L., Glasius, M., and Bilde, M.: The effect of sub-zero temperature on the formation and composition of secondary organic aerosol from ozonolysis of alpha-pinene, Environmental Science: Processes & Impacts, 19, 1220-1234, 2017.

Kuang, C., McMurry, P. H., McCormick, A. V., and Eisele, F. L.: Dependence of nucleation rates on sulfuric acid vapor concentration in diverse atmospheric locations, Journal of Geophysical Research: Atmospheres, 113, 2008.

900   Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T., Petaja, T., Sipila, M., Schobesberger, S., Rantala, P., Franchin, A., Jokinen, T., Jarvinen, E., Aijala, M., Kangasluoma, J., Hakala, J., Aalto, P. P., Paasonen, P., Mikkila, J., Vanhanen, J., Aalto, J., Hakola, H., Makkonen, U., Ruuskanen, T., Mauldin, R. L., Duplissy, J., Vehkamaki, H., Back, J., Kortelainen, A., Riipinen, I., Kurten, T., Johnston, M. V., Smith, J. N., Ehn, M., Mentel, T. F., Lehtinen, K. E. J., Laaksonen, A., Kerminen, V. M., and Worsnop, D. R.: Direct Observations of Atmospheric Aerosol Nucleation,

905   Science, 339, 943-946, 2013.

Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V. M., Birmili, W., and McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: a review of observations, Journal of Aerosol Science, 35, 143-176, 2004.

Kupc, A., Amorim, A., Curtius, J., Danielczok, A., Duplissy, J., Ehrhart, S., Walther, H., Ickes, L., Kirkby, J., and Kürten, A.:

910   A fibre-optic UV system for H2SO4 production in aerosol chambers causing minimal thermal effects, Journal of Aerosol Science, 42, 532-543, 2011.

Kürten, A., Bianchi, F., Almeida, J., Kupiainen-Määttä, O., Dunne, E. M., Duplissy, J., Williamson, C., Barmet, P., Breitenlechner, M., and Dommen, J.: Experimental particle formation rates spanning tropospheric sulfuric acid and

ammonia abundances, ion production rates, and temperatures, Journal of Geophysical Research: Atmospheres, 121, 12,377-312,400, 2016.

Kürten, A., Jokinen, T., Simon, M., Sipilä, M., Sarnela, N., Junninen, H., Adamov, A., Almeida, J., Amorim, A., Bianchi, F., Breitenlechner, M., Dommen, J., Donahue, N. M., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Hakala, J., Hansel, A., Heinritzi, M., Hutterli, M., Kangasluoma, J., Kirkby, J., Laaksonen, A., Lehtipalo, K., Leiminger, M., Makhmutov, V., Mathot, S., Onnela, A., Petäjä, T., Praplan, A. P., Riccobono, F., Rissanen, M. P., Rondo, L., Schobesberger, S., Seinfeld, J. H., Steiner, G., Tomé, A., Tröstl, J., Winkler, P. M., Williamson, C., Wimmer, D., Ye, P., Baltensperger, U., Carslaw, K. S., Kulmala, M., Worsnop, D. R., and Curtius, J.: Neutral molecular cluster formation of sulfuric acid–dimethylamine observed in real time under atmospheric conditions, Proceedings of the National Academy of Sciences, 111, 15019-15024, 2014.

Kürten, A., Rondo, L., Curtius, J., and Ehrhart, S.: Performance of a corona ion source for measurement of sulfuric acid by chemical ionization mass spectrometry, Atmos. Meas. Tech., 4, 437-443, 2011.

Kürten, A., Rondo, L., Ehrhart, S., and Curtius, J.: Calibration of a Chemical Ionization Mass Spectrometer for the Measurement of Gaseous Sulfuric Acid, The Journal of Physical Chemistry A, 116, 6375-6386, 2012.

Kürten, A., Williamson, C., Almeida, J., Kirkby, J., and Curtius, J.: On the derivation of particle nucleation rates from experimental formation rates, Atmos. Chem. Phys., 15, 4063-4075, 2015.

Kurtén, T., Loukonen, V., Vehkamäki, H., and Kulmala, M.: Amines are likely to enhance neutral and ion-induced sulfuric acid-water nucleation in the atmosphere more effectively than ammonia, Atmospheric Chemistry and Physics, 8, 4095-4103, 2008.

Kurtén, T., Rissanen, M. P., Mackeprang, K., Thornton, J. A., Hyttinen, N., Jørgensen, S., Ehn, M., and Kjaergaard, H. G.: Computational study of hydrogen shifts and ring-opening mechanisms in α-pinene ozonolysis products, The Journal of Physical Chemistry A, 119, 11366-11375, 2015.

Kurtén, T., Tiusanen, K., Roldin, P., Rissanen, M., Luy, J.-N., Boy, M., Ehn, M., and Donahue, N.: α-Pinene Autoxidation Products May Not Have Extremely Low Saturation Vapor Pressures Despite High O:C Ratios, The Journal of Physical Chemistry A, 120, 2569-2582, 2016.

Lehtipalo, K., Leppä, J., Kontkanen, J., Kangasluoma, J., Franchin, A., Wimmer, D., Schobesberger, S., Junninen, H., Petäjä, T., Sipilä, M., Mikkilä, J., Vanhanen, J., Worsnop, D. R., and Kulmala, M.: Methods for determining particle size distribution and growth rates between 1 and 3 nm using the Particle Size Magnifier, Boreal Env. Res., 19 (suppl. B), 215-236, 2014.

Lehtipalo, K., Yan, C., Dada, L., Bianchi, F., Xiao, M., Wagner, R., Stolzenburg, D., Ahonen, L. R., Amorim, A., and Baccarini, A.: Multicomponent new particle formation from sulfuric acid, ammonia, and biogenic vapors, Science Advances, 4, eaau5363, 2018.

Lopez-Hilfiker, F. D., Mohr, C., Ehn, M., Rubach, F., Kleist, E., Wildt, J., Mentel, T. F., Lutz, A., Hallquist, M., Worsnop, D., and Thornton, J. A.: A novel method for online analysis of gas and particle composition: description and evaluation of a Filter Inlet for Gases and AEROsols (FIGAERO), Atmos. Meas. Tech., 7, 983-1001, 2014.

Manninen, H. E., Petäjä, T., Asmi, E., Riipinen, I., Nieminen, T., Mikkilä, J., Hõrrak, U., Mirme, A., Mirme, S., and Laakso, L.: Long-term field measurements of charged and neutral clusters using Neutral cluster and Air Ion Spectrometer (NAIS), 2009. 2009.

Merikanto, J., Spracklen, D. V., Mann, G. W., Pickering, S. J., and Carslaw, K. S.: Impact of nucleation on global CCN, Atmos. Chem. Phys., 9, 8601-8616, 2009.

Mohr, C., Lopez-Hilfiker, F. D., Yli-Juuti, T., Heitto, A., Lutz, A., Hallquist, M., D'Ambro, E. L., Rissanen, M. P., Hao, L., and Schobesberger, S.: Ambient observations of dimers from terpene oxidation in the gas phase: Implications for new particle formation and growth, Geophysical Research Letters, 44, 2958-2966, 2017.

Molteni, U., Simon, M., Heinritzi, M., Hoyle, C. R., Bernhammer, A.-K., Bianchi, F., Breitenlechner, M., Brilke, S., Dias, A., and Duplissy, J.: Formation of Highly Oxygenated Organic Molecules from alpha-Pinene Ozonolysis, ACS Earth and Space Chemistry, 2019. 2019.

Ng, N., Chhabra, P., Chan, A., Surratt, J., Kroll, J., Kwan, A., McCabe, D., Wennberg, P., Sorooshian, A., and Murphy, S.: Effect of NO x level on secondary organic aerosol (SOA) formation from the photooxidation of terpenes, Atmospheric Chemistry and Physics, 7, 5159-5174, 2007.

O'Dowd, C. D., Jimenez, J. L., Bahreini, R., Flagan, R. C., Seinfeld, J. H., Hämeri, K., Pirjola, L., Kulmala, M., Jennings, S. G., and Hoffmann, T.: Marine aerosol formation from biogenic iodine emissions, Nature, 417, 632, 2002.

Pankow, J. F. and Asher, W. E.: SIMPOL. 1: a simple group contribution method for predicting vapor pressures and enthalpies of vaporization of multifunctional organic compounds, Atmospheric Chemistry and Physics, 8, 2773-2796, 2008.

Passananti, M., Zapadinsky, E., Zanca, T., Kangasluoma, J., Myllys, N., Rissanen, M. P., Kurtén, T., Ehn, M., Attoui, M., and Vehkamäki, H.: How well can we predict cluster fragmentation inside a mass spectrometer?, Chemical Communications, 55, 5946-5949, 2019.

Praske, E., Otkjær, R. V., Crounse, J. D., Hethcox, J. C., Stoltz, B. M., Kjaergaard, H. G., and Wennberg, P. O.: Atmospheric autoxidation is increasingly important in urban and suburban North America, Proceedings of the National Academy of Sciences, 115, 64-69, 2018.

Presto, A. A., Huff Hartz, K. E., and Donahue, N. M.: Secondary organic aerosol production from terpene ozonolysis. 2. Effect of NO x concentration, Environmental Science & Technology, 39, 7046-7054, 2005.

975 Quéléver, L. L., Kristensen, K., Normann Jensen, L., Rosati, B., Teiwes, R., Daellenbach, K. R., Peräkylä, O., Roldin, P., Bossi, R., and Pedersen, H. B.: Effect of temperature on the formation of highly oxygenated organic molecules (HOMs) from alpha-pinene ozonolysis, Atmospheric Chemistry and Physics, 19, 7609-7625, 2019.

Riccobono, F., Schobesberger, S., Scott, C. E., Dommen, J., Ortega, I. K., Rondo, L., Almeida, J., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Downard, A., Dunne, E. M., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Hansel, A., Junninen, H., Kajos, M., Keskinen, H., Kupc, A., Kürten, A., Kvashin, A. N., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Petäjä, T., Praplan, A. P., Santos, F. D., Schallhart, S., Seinfeld, J. H., Sipilä, M., Spracklen, D. V., Stozhkov, Y., Stratmann, F., Tomé, A., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y., Vrtala, A., Wagner, P. E., Weingartner, E., Wex, H., Wimmer, D., Carslaw, K. S., Curtius, J., Donahue, N. M., Kirkby, J., Kulmala, M., Worsnop, D. R., and Baltensperger, U.: Oxidation Products of Biogenic Emissions Contribute to Nucleation of Atmospheric Particles, Science, 344, 717, 2014.

Riipinen, I., Yli-Juuti, T., Pierce, J. R., Petäjä, T., Worsnop, D. R., Kulmala, M., and Donahue, N. M.: The contribution of organics to atmospheric nanoparticle growth, Nature Geoscience, 5, 453-458, 2012.

Rissanen, M. P.: NO2 Suppression of Autoxidation–Inhibition of Gas-Phase Highly Oxidized Dimer Product Formation, ACS Earth and Space Chemistry, 2, 1211-1219, 2018.

990 Rissanen, M. P., Kurtén, T., Sipilä, M., Thornton, J. A., Kangasluoma, J., Sarnela, N., Junninen, H., Jørgensen, S., Schallhart, S., and Kajos, M. K.: The formation of highly oxidized multifunctional products in the ozonolysis of cyclohexene, Journal of the American Chemical Society, 136, 15596-15606, 2014.

Rissanen, M. P., Kurtén, T., Sipilä, M., Thornton, J. A., Kausiala, O., Garmash, O., Kjaergaard, H. G., Petäjä, T., Worsnop, D. R., Ehn, M., and Kulmala, M.: Effects of Chemical Complexity on the Autoxidation Mechanisms of Endocyclic Alkene Ozonolysis Products: From Methylcyclohexenes toward Understanding α-Pinene, The Journal of Physical Chemistry A, 119, 4633-4650, 2015.

Saathoff, H., Naumann, K. H., Möhler, O., Jonsson, Å. M., Hallquist, M., Kiendler-Scharr, A., Mentel, T. F., Tillmann, R., and Schurath, U.: Temperature dependence of yields of secondary organic aerosols from the ozonolysis of <i>α</i>-pinene and limonene, Atmos. Chem. Phys., 9, 1551-1577, 2009.

1000 Schervish, M. and Donahue, N. M.: Peroxy Radical Chemistry and the Volatility Basis Set, Atmos. Chem. Phys. Discuss., 2019, 1-29, 2019.

Schnitzhofer, R., Metzger, A., Breitenlechner, M., Jud, W., Heinritzi, M., De Menezes, L. P., Duplissy, J., Guida, R., Haider, S., Kirkby, J., Mathot, S., Minginette, P., Onnela, A., Walther, H., Wasem, A., Hansel, A., and the, C. T.: Characterisation of organic contaminants in the CLOUD chamber at CERN, Atmos. Meas. Tech., 7, 2159-2168, 2014.

1005 Schobesberger, S., D'Ambro, E. L., Lopez-Hilfiker, F. D., Mohr, C., and Thornton, J. A.: A model framework to retrieve thermodynamic and kinetic properties of organic aerosol from composition-resolved thermal desorption measurements, Atmos. Chem. Phys., 18, 14757-14785, 2018.

Schobesberger, S., Junninen, H., Bianchi, F., Lönn, G., Ehn, M., Lehtipalo, K., Dommen, J., Ehrhart, S., Ortega, I. K., Franchin, A., Nieminen, T., Riccobono, F., Hutterli, M., Duplissy, J., Almeida, J., Amorim, A., Breitenlechner, M., Downard, A. J.,
1010 Dunne, E. M., Flagan, R. C., Kajos, M., Keskinen, H., Kirkby, J., Kupc, A., Kürten, A., Kurtén, T., Laaksonen, A., Mathot, S., Onnela, A., Praplan, A. P., Rondo, L., Santos, F. D., Schallhart, S., Schnitzhofer, R., Sipilä, M., Tomé, A., Tsagkogeorgas, G., Vehkamäki, H., Wimmer, D., Baltensperger, U., Carslaw, K. S., Curtius, J., Hansel, A., Petäjä, T., Kulmala, M., Donahue, N. M., and Worsnop, D. R.: Molecular understanding of atmospheric particle formation from sulfuric acid and large oxidized organic molecules, Proceedings of the National Academy of Sciences, 110, 17223-17228,
1015 2013.

Simon, M., Heinritzi, M., Herzog, S., Leiminger, M., Bianchi, F., Praplan, A., Dommen, J., Curtius, J., and Kürten, A.: Detection of dimethylamine in the low pptv range using nitrate chemical ionization atmospheric pressure interface time-of-flight (CI-APi-TOF) mass spectrometry, Atmospheric Measurement Techniques, 9, 2135-2145, 2016.

Sindelarova, K., Granier, C., Bouarar, I., Guenther, A., Tilmes, S., Stavrakou, T., Müller, J. F., Kuhn, U., Stefani, P., and
1020 Knorr, W.: Global data set of biogenic VOC emissions calculated by the MEGAN model over the last 30 years, Atmos. Chem. Phys., 14, 9317-9341, 2014.

Sipilä, M., Sarnela, N., Jokinen, T., Henschel, H., Junninen, H., Kontkanen, J., Richters, S., Kangasluoma, J., Franchin, A., Peräkylä, O., Rissanen, M. P., Ehn, M., Vehkamäki, H., Kurten, T., Berndt, T., Petäjä, T., Worsnop, D., Ceburnis, D., Kerminen, V.-M., Kulmala, M., and O'Dowd, C.: Molecular-scale evidence of aerosol particle formation via sequential
1025 addition of HIO3, Nature, 537, 532, 2016.

Spracklen, D. V., Carslaw, K. S., Kulmala, M., Kerminen, V. M., Sihto, S. L., Riipinen, I., Merikanto, J., Mann, G. W., Chipperfield, M. P., and Wiedensohler, A.: Contribution of particle formation to global cloud condensation nuclei concentrations, Geophysical Research Letters, 35, 2008.

Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P.
1030 M.: Climate change 2013: The physical science basis, Intergovernmental Panel on Climate Change, Working Group I Contribution to the IPCC Fifth Assessment Report (AR5)(Cambridge Univ Press, New York), 25, 2013.

Stolzenburg, D., Fischer, L., Vogel, A. L., Heinritzi, M., Schervish, M., Simon, M., Wagner, A. C., Dada, L., Ahonen, L. R., and Amorim, A.: Rapid growth of organic aerosol nanoparticles over a wide tropospheric temperature range, Proceedings of the National Academy of Sciences, 2018. 201807604, 2018.

1035 Stolzenburg, D., Steiner, G., and Winkler, P. M.: A DMA-train for precision measurement of sub-10 nm aerosol dynamics, Atmos. Meas. Tech., 10, 1639-1651, 2017.

Tröstl, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Ahlm, L., Frege, C., Bianchi, F., Wagner, R., Simon, M., Lehtipalo, K., Williamson, C., Craven, J. S., Duplissy, J., Adamov, A., Almeida, J., Bernhammer, A.-K., Breitenlechner, M., Brilke, S., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Gysel, M., Hansel, 1040 A., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Keskinen, H., Kim, J., Krapf, M., Kürten, A., Laaksonen, A., Lawler, M., Leiminger, M., Mathot, S., Möhler, O., Nieminen, T., Onnela, A., Petäjä, T., Piel, F. M., Miettinen, P., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Sengupta, K., Sipilä, M., Smith, J. N., Steiner, G., Tomè, A., Virtanen, A., Wagner, A. C., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Carslaw, K. S., Curtius, J., Dommen, J., Kirkby, J., Kulmala, M., Riipinen, I., Worsnop, D. R., Donahue, N. M., and Baltensperger, U.: The role of low-volatility 1045 organic compounds in initial particle growth in the atmosphere, Nature, 533, 527, 2016.

Valiev, R. R., Hasan, G., Salo, V.-T., Kubečka, J., and Kurten, T.: Intersystem Crossings Drive Atmospheric Gas-Phase Dimer Formation, The Journal of Physical Chemistry A, 123, 6596-6604, 2019.

Vanhanen, J., Mikkilä, J., Lehtipalo, K., Sipilä, M., Manninen, H. E., Siivola, E., Petäjä, T., and Kulmala, M.: Particle Size Magnifier for Nano-CN Detection, Aerosol Science and Technology, 45, 533-542, 2011.

1050 Viggiano, A., Seeley, J. V., Mundis, P. L., Williamson, J. S., and Morris, R. A.: Rate Constants for the Reactions of XO3-(H2O) n (X= C, HC, and N) and NO3-(HNO3) n with H2SO4: Implications for Atmospheric Detection of H2SO4, The Journal of Physical Chemistry A, 101, 8275-8278, 1997.

Wagner, R., Yan, C., Lehtipalo, K., Duplissy, J., Nieminen, T., Kangasluoma, J., Ahonen, L. R., Dada, L., Kontkanen, J., and Manninen, H. E.: The role of ions in new particle formation in the CLOUD chamber, Atmospheric Chemistry and Physics, 1055 17, 15181-15197, 2017.

Werle, P., Slemr, F., Maurer, K., Kormann, R., Mücke, R., and Jänker, B.: Near-and mid-infrared laser-optical sensors for gas analysis, Optics and lasers in engineering, 37, 101-114, 2002.

WHO: Ambient air pollution: A global assessment of exposure and burden of disease, 2016. 2016.

Wildt, J., Mentel, T., Kiendler-Scharr, A., Hoffmann, T., Andres, S., Ehn, M., Kleist, E., Müsgen, P., Rohrer, F., and Rudich, 1060 Y.: Suppression of new particle formation from monoterpene oxidation by NO x, Atmospheric chemistry and physics, 14, 2789-2804, 2014.

Williamson, C. J., Kupc, A., Axisa, D., Bilsback, K. R., Bui, T., Campuzano-Jost, P., Dollner, M., Froyd, K. D., Hodshire, A. L., Jimenez, J. L., Kodros, J. K., Luo, G., Murphy, D. M., Nault, B. A., Ray, E. A., Weinzierl, B., Wilson, J. C., Yu, F., Yu, P., Pierce, J. R., and Brock, C. A.: A large source of cloud condensation nuclei from new particle formation in the tropics, Nature, 574, 399-403, 2019.

Wimmer, D., Lehtipalo, K., Franchin, A., Kangasluoma, J., Kreissl, F., Kürten, A., Kupc, A., Metzger, A., Mikkilä, J., and Petäjä, T.: Performance of diethylene glycol-based particle counters in the sub-3 nm size range, Atmospheric Measurement Techniques, 6, 1793-1804, 2013.

Wimmer, D., Lehtipalo, K., Nieminen, T., Duplissy, J., Ehrhart, S., Almeida, J., Rondo, L., Franchin, A., Kreissl, F., Bianchi, F., Manninen, H. E., Kulmala, M., Curtius, J., and Petäjä, T.: Technical Note: Using DEG-CPCs at upper tropospheric temperatures, Atmos. Chem. Phys., 15, 7547-7555, 2015.

Ye, P., Zhao, Y., Chuang, W. K., Robinson, A. L., and Donahue, N. M.: Secondary organic aerosol production from pinanediol, a semi-volatile surrogate for first-generation oxidation products of monoterpenes, Atmos. Chem. Phys., 18, 6171-6186, 2018.

Ye, Q., Wang, M., Hofbauer, V., Stolzenburg, D., Chen, D., Schervish, M., Vogel, A. L., Mauldin, R. L., Baalbaki, R., Brilke, S., Dada, L., Dias, A., Duplissy, J., El Haddad, I., Finkenzeller, H., Fischer, L., He, X., Kim, C., Kürten, A., Lamkaddam, H., Lee, C. P., Lehtipalo, K., Leiminger, M., Manninen, H. E., Marten, R., Mentler, B., Partoll, E., Petäjä, T., Rissanen, M. P., Schobesberger, S., Schuchmann, S., Simon, M., Tham, Y. J., Vazquez-Pufleau, M., Wagner, A. C., Wang, Y., Wu, Y., Xiao, M., Baltensperger, U., Curtius, J., Flagan, R., Kirkby, j., Kulmala, M., Volkamer, R., Winkler, P. M., Worsnop, D. R., and Donahue, N. M.: Molecular Composition and Volatility of Nucleated Particles from α-Pinene Oxidation between -50 °C and +25 °C, Environmental Science & Technology, doi: 10.1021/acs.est.9b03265, 2019. 2019.

Ziemann, P. J. and Atkinson, R.: Kinetics, products, and mechanisms of secondary organic aerosol formation, Chemical Society Reviews, 41, 6582-6605, 2012.

1085
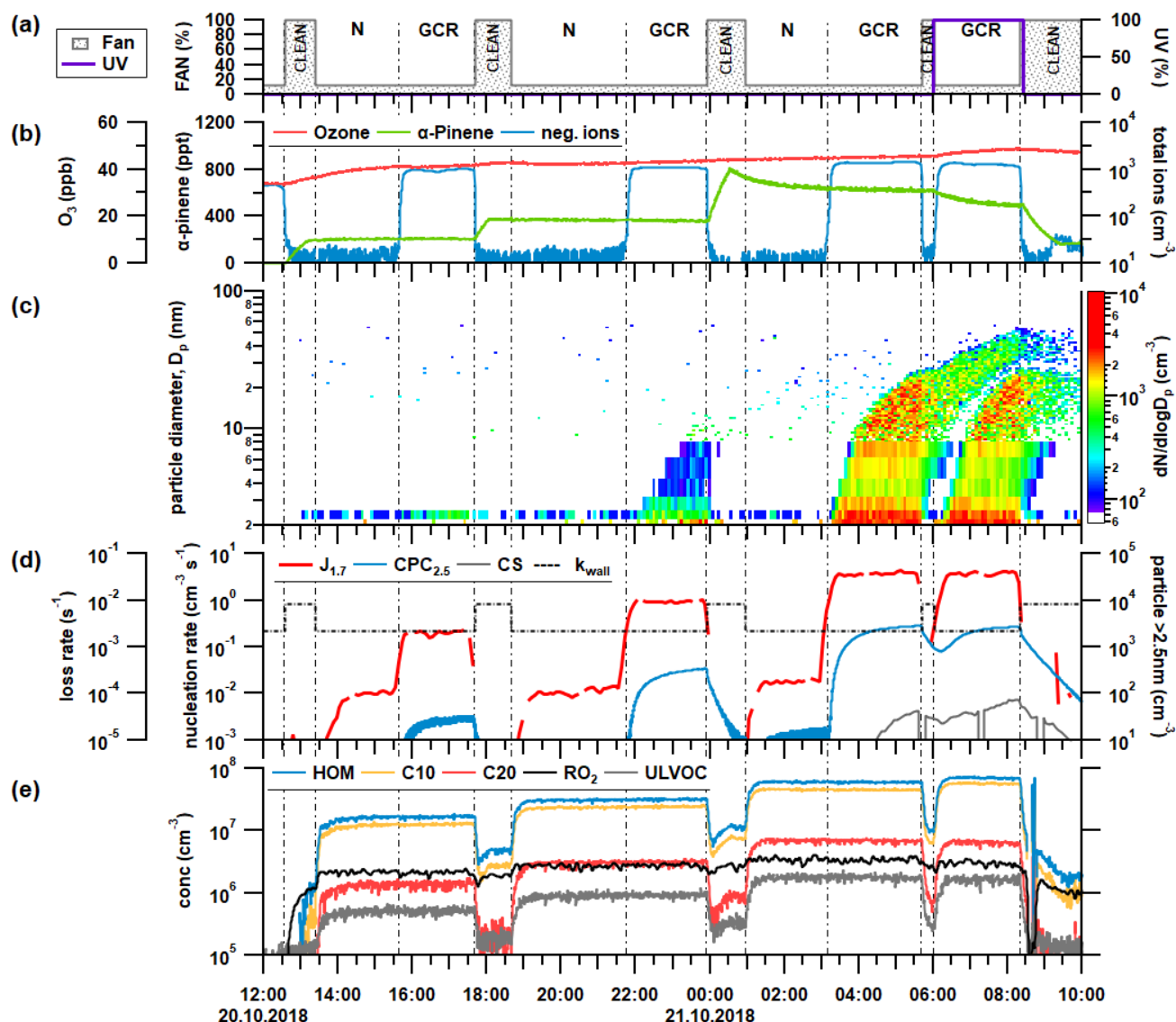
Figure 1: Typical CLOUD experiment sequence of a new-particle formation experiment by α-pinene oxidation for three different precursor concentrations.

The figure shows an example α-pinene new-particle formation run during the CLOUD13 campaign. The experiment is conducted at a temperature of + 5 °C and at a relative humidity of 80 %. The vertical lines indicate a change of the experimental

1090 conditions in the chamber (e.g. change of settings for fan speed, UV illumination, clearing field, etc.) marking a new stage within the run. (A) Change of fan speed and UV light intensity during the run. N, GCR and CLEAN indicate neutral (high-voltage clearing field on), galactic cosmic ray (high-voltage clearing field off) and cleaning (neutral periods including high fan speed to clean the chamber of particles) conditions, respectively. (B) Time series of ozone, α-pinene and negative ions. (C)

Combined size distribution of aerosol particles measured by the DMA-TRAIN (1.8–8 nm) and nanoSMPS (8–63 nm). (D)
Evolution of the nucleation rate at 1.7nm ($J_{1.7}$) and the total particle concentration above 2.5 nm, measured with a scanning
PSM (1.7 nm) and a butanol-based CPC (2.5 nm). Furthermore, the loss rates to the chamber walls (k_{wall}, black dashed line)
and the determined particle condensation sink (CS, grey line) are shown. (E) Evolution of total HOM concentration and
partitioning into HOM monomer (C_{10}), HOM dimer (C_{20}) and peroxy radicals (RO_2) as well the fraction of ultra-low volatility
organic compounds (ULVOC) measured by the nitrate CI-APi-TOF.
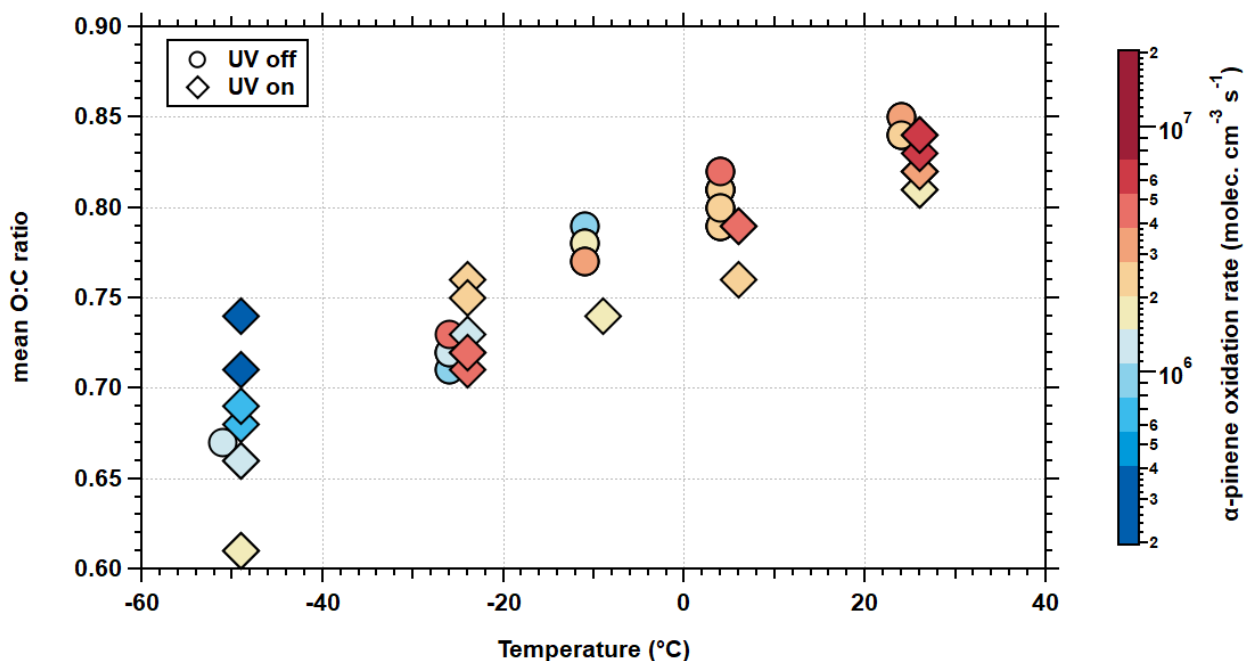
Figure 2: Mean oxygen-to-carbon elemental ratio (O/C) of HOMs at different temperatures and various α-pinene oxidation rates.

Mean O/C of HOMs detected by the nitrate CI-APi-TOF. HOMs are produced from a mixture of α-pinene (200–2000 pptᵥ)
and ozone (37–48 ppbᵥ) under dark (circles) and UV-illuminated (diamonds) conditions at five different temperatures ranging
from - 50 to + 25 °C. The marker symbols are slightly horizontally shifted on the x-axis to ensure a better visibility. The higher
OH radical concentration due to the UV light influences the oxidation state of the HOM molecules. The different initial
precursor concentrations lead to only a minor effect on the oxidation ratio. Temperature has an impact on the formation of
HOM closed-shell products as the autoxidation proceeds slower at lower temperatures.
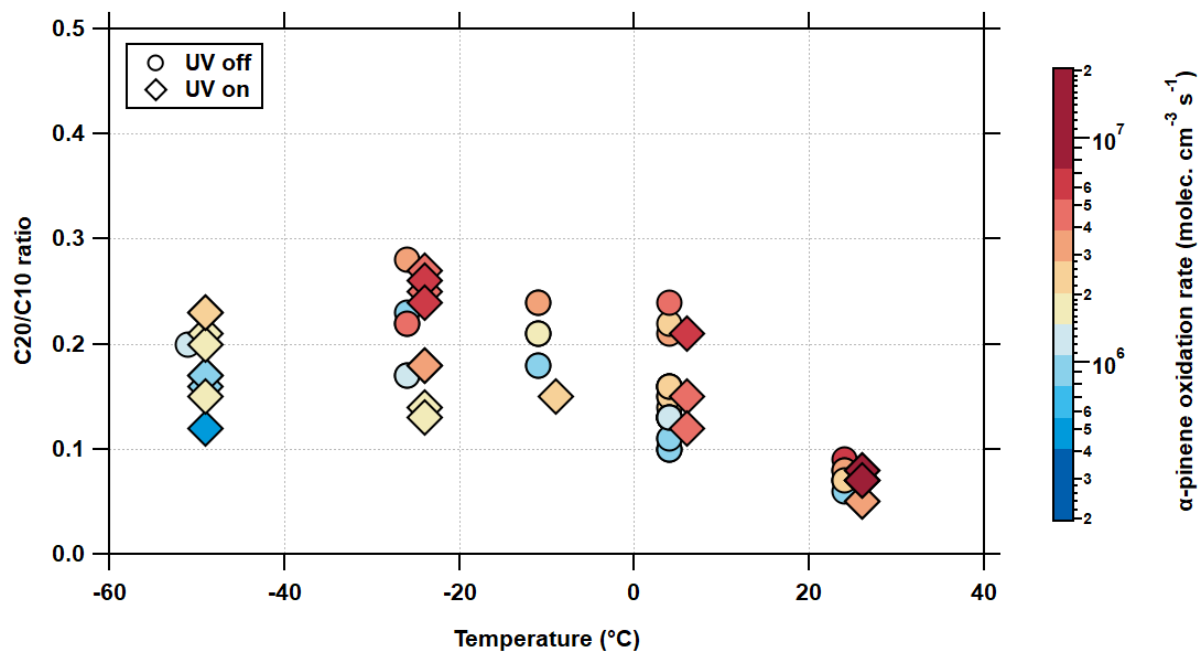
Figure 3: Mean C$_{20}$ dimer to C$_{10}$ monomer ratio of α-pinene oxidation at different temperatures.

The ratio of the detected sum of dimer products (C$_{20}$) versus the sum of monomer products of the α-pinene oxidation experiment measured by the nitrate CI-APi-TOF. The color bar indicates the α-pinene oxidation rate. Elevated oxidation rates result in a stronger increase in dimer products than monomer products due to more RO$_2$-RO$_2$ reactions. While the ratio between monomers and dimers is stable for low temperatures (< 0 °C) it decreases with higher temperatures. This is an indication that some of the monomers are semi-volatile at higher temperatures and do not irreversibly condense on a cluster or particle. Also, the bimolecular termination of the autoxidation, due to an elevated concentration of OH radicals, can cause this effect. No pronounced effect due to the different ion concentration in the chamber experiments can be recognized.

1125
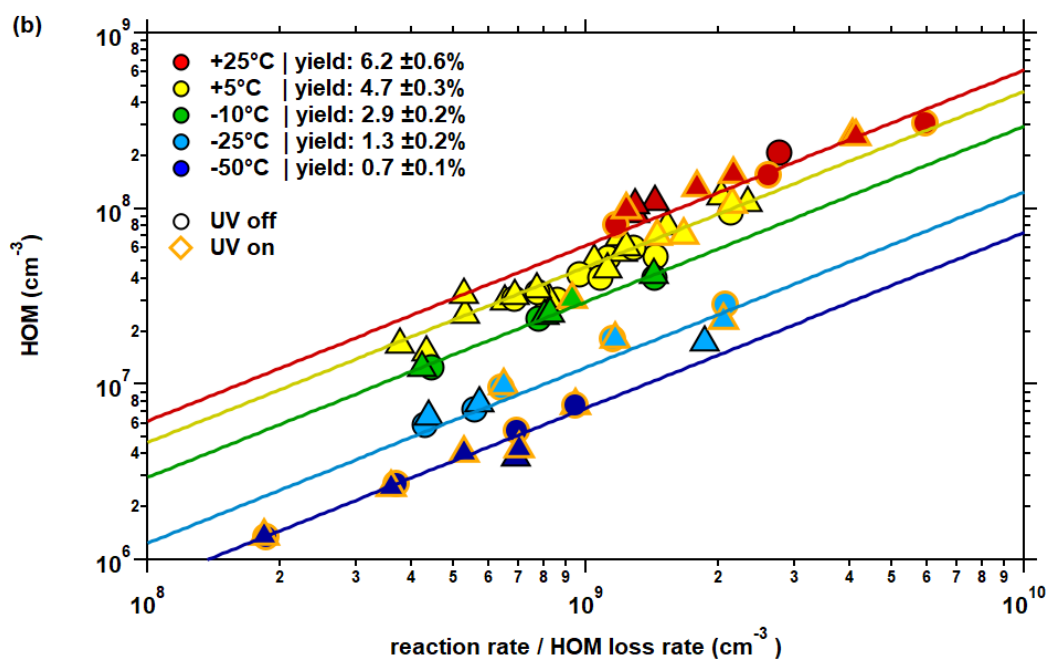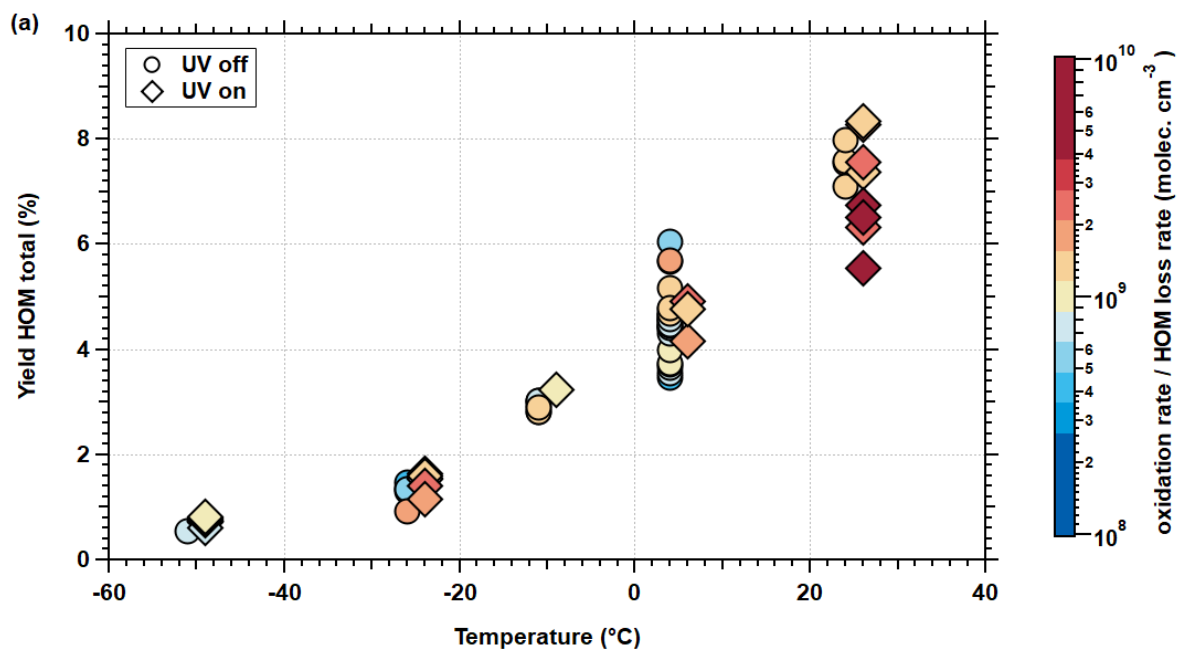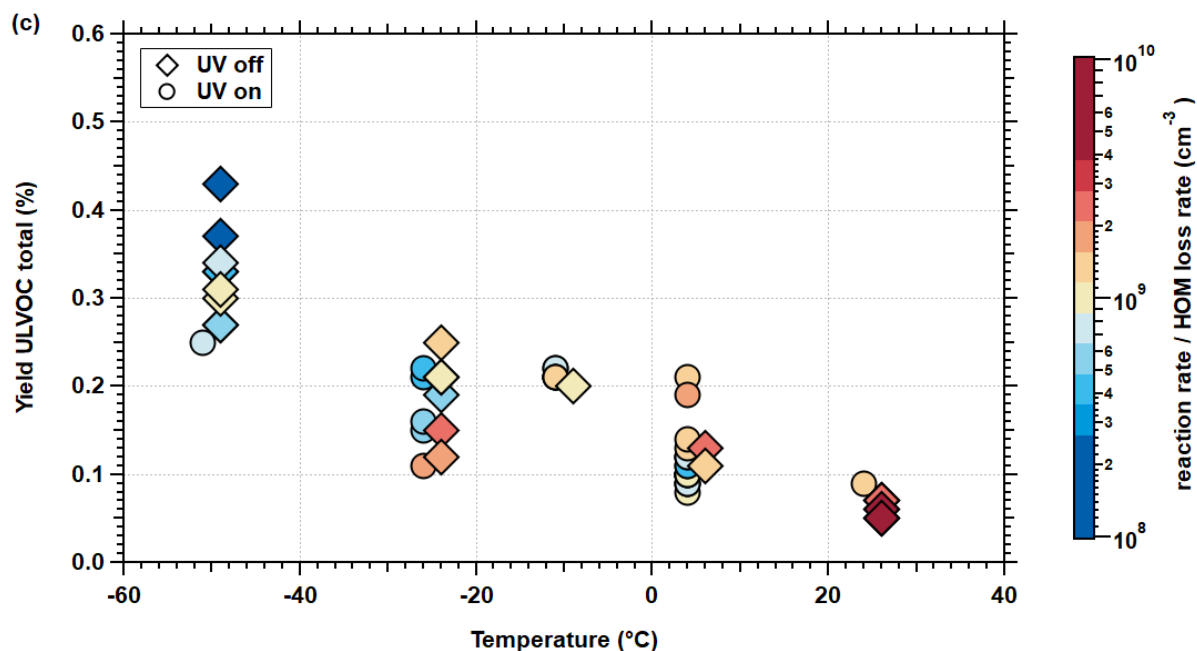
Figure 4: Total HOM and ULVOC yields as a function of α-pinene oxidation rates at five different temperatures.

The total gas-phase (a) HOM and (c) ULVOC yields measured by the nitrate CI-APi-TOF according to Eq. (10). The color code in panel (a) and (c) as well as the x axis in panel B shows the oxidation rate corrected for all losses and condensation sink inside the chamber. (a) The total HOM yield decreases with decreasing temperature. The marker symbols for dark (circles) and UV-illuminated (diamonds) conditions are slightly horizontally shifted on the x-axis to ensure a better visibility. (b) The mean yields are derived from the slope of the linear fits of the HOM concentrations at the different temperatures. The measured HOM concentration is linearly dependent on its precursor concentration (oxidation rate) as indicated by the double-logarithmic fits with the slope of 1. Thus, there is no saturation effect in the chemical production rate nor an instrumental effect at all temperatures. Overall systematic scale uncertainties of ± 40 % for the α-pinene reaction rates and + 78%/− 68% for the HOM mixing ratios are not shown.
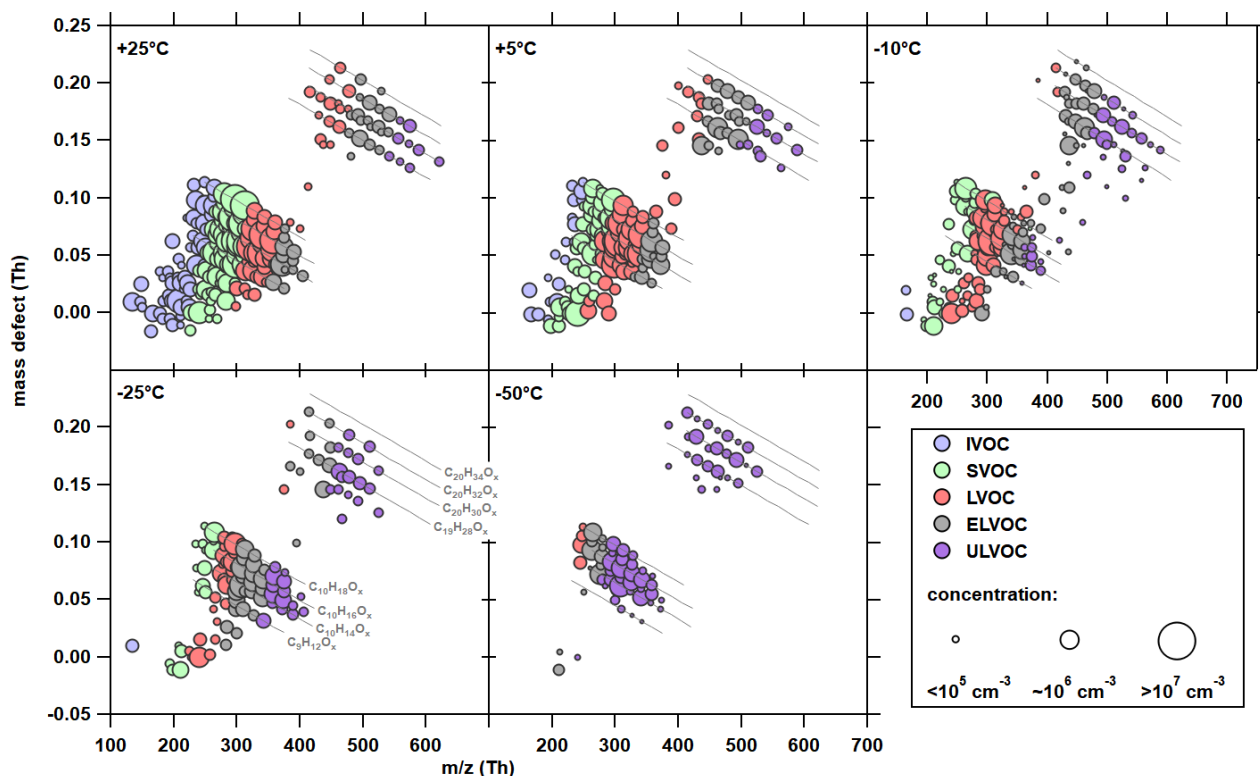
Figure 5: Molecular composition of neutral highly-oxygenated molecules (HOM) at the five temperatures studied and their corresponding volatility.

Mass defect (difference of exact mass to integer mass) versus m/z of gas-phase HOMs measured with the nitrate anion CI-APi-TOF. Data shown represent steady-state run conditions (production term balanced with all loss rate). For clarity, other signals, such as the reagent ions, are not shown in the plot. Each circle represents a particular molecular composition. All molecules shown here are clustered to a nitrate ion monomer from the reagent ion chemical ionization. Same oxidation products clustered to different reagent ions (e.g. $(HNO_3)NO_3^- \cdot C_{10}H_{14}O_7$) are summed to the reagent monomer cluster ion ($NO_3^- \cdot C_{10}H_{14}O_7$). Similar α-pinene oxidation rates ($2 - 3 \cdot 10^6$ molecule $\cdot$ cm^{-3} $\cdot$ s^{-1}) are used for each temperature to ensure better comparability. The area of the markers is proportional to the logarithm of the concentration measured by the nitrate CI-APi-TOF. The systematic scale uncertainty on the HOM concentrations is $+ 78$ %/$- 68$ %. The lines represent a homologous group of oxidation products with the same hydrogen to carbon (H/C) elemental ratio but with an increasing oxygen level (e.g. $C_{10}H_{14}O_{3-12}$). The group between 240 and 420 Th is referred to as HOM monomers (mainly C_8 - C_{10}), while the band between 400 and 620 Th marks HOM dimers (mainly C_{18}–C_{20}). The color indicates the corresponding volatility range based on the volatility basis set model (VBS-model). ULVOC, ultra-low volatility organic compound (purple; $C^*(300K) < 3 \cdot 10^{-9}$ µg m^{-3}); ELVOC, extremely low-volatility organic compound (gray; $3 \cdot 10^{-9} < C^*(300K) < 3 \cdot 10^{-5}$ µg m^{-3}); LVOC, low-volatility organic compound (red, $3 \cdot 10^{-5} < C^*(300K) < 0.3$ µg m^{-3}); SVOC, semi-volatile organic compound

(green, $0.3 < C^*(300K) < 300 \, \mu g \, m^{-3}$); IVOC, intermediate volatility organic compound (blue, $300 < C^*(300K) < 3 \cdot 10^6 \, \mu g \, m^{-3}$).
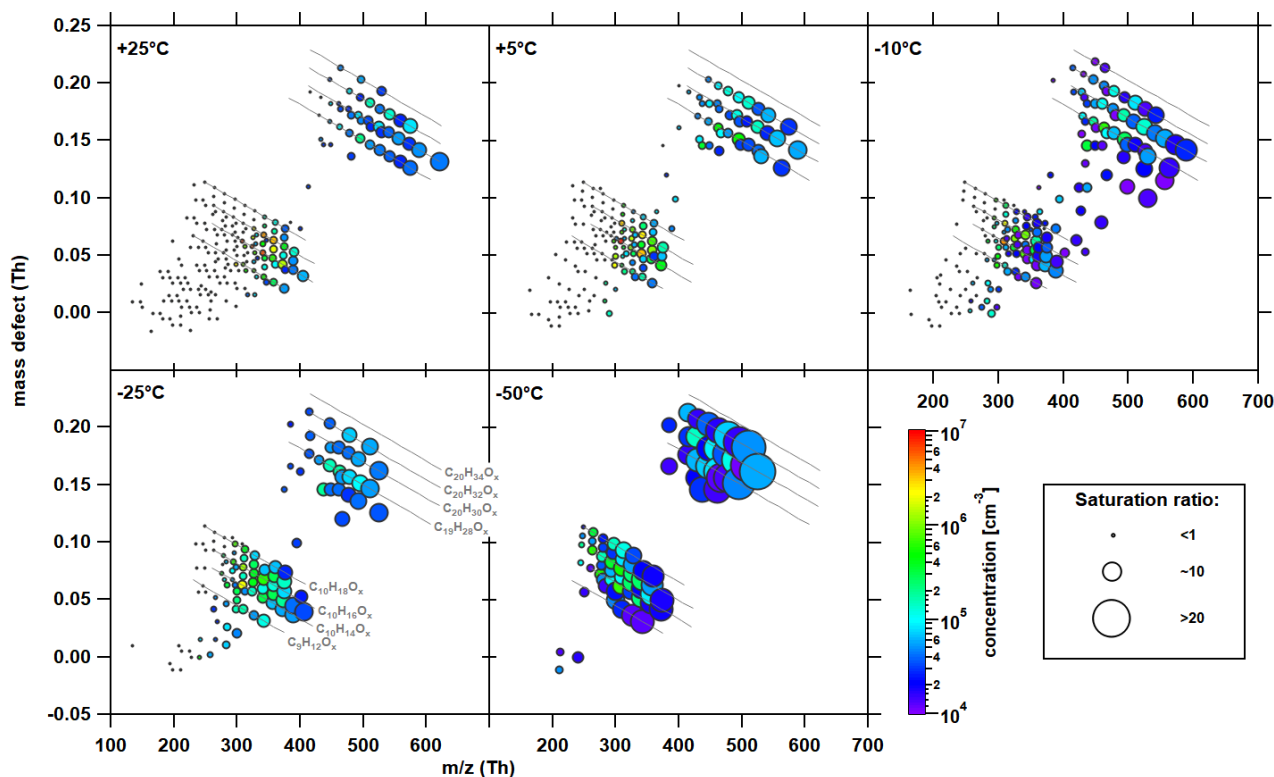
Figure 6: Composition of the measured gas-phase HOMs at different temperatures and their saturation ratio distribution.

Same α-pinene oxidation rates ($2 - 3 \cdot 10^6$ molecule $\cdot$ cm^{-3} $\cdot$ s^{-1}) and run conditions as shown in Fig. 5. The area of the markers is proportional to the saturation ratio of each detected HOM in the gas-phase as measured by the nitrate CI-APi-TOF, while the color reflects the measured concentration. The figure directly shows the molecules that drive the nucleation and early growth of embryonic clusters. Note that the concentrations of less oxidized products are likely underestimated by the nitrate CI-APi-TOF and should therefore be interpreted as a lower limit. The lines represent a homologous group of oxidation products with the same H/C but different number of oxygen atoms.
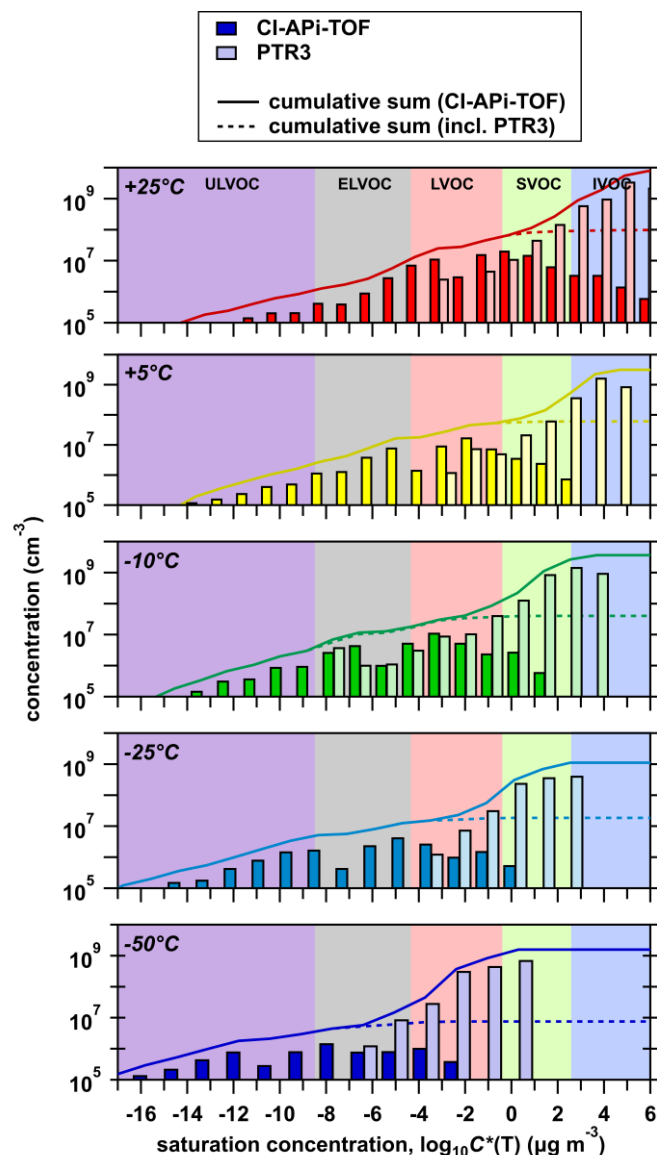
Figure 7: Volatility distribution of the measured oxidation products for five representative temperatures.

1175 The mass concentration and volatility distributions of HOMs shown here are at the same α-pinene oxidation rates $(2-3 \cdot 10^6 \text{ molecule} \cdot \text{cm}^{-3} \cdot \text{s}^{-1})$ and run conditions as data shown in Fig. 5. The summed HOM concentration of each volatility bin is divided into the nitrate CI-APi-TOF (dark color) and the PTR3 instrument (light color), respectively. Each volatility bin is defined at 300K, shifted and widened according to their corresponding temperature. The highest and lowest bin are overflow bins. The solid lines for each temperature represent the cumulative sum of the volatility bins of both instruments towards

 higher volatility, while the dashed line represents the nitrate CIMS only. HOMs detected by the nitrate CI-APi-TOF (LOD ≤ 5·10⁴ molecules·cm⁻³) below 1·10⁵ molecules·cm⁻³ are added to the cumulative sum. This results in slightly higher concentrations of the cumulative sums when compared with the first visible bin. The color bands in the background indicate the volatility regimes as in Donahue et al. (2012) and Schervish and Donahue (2019).
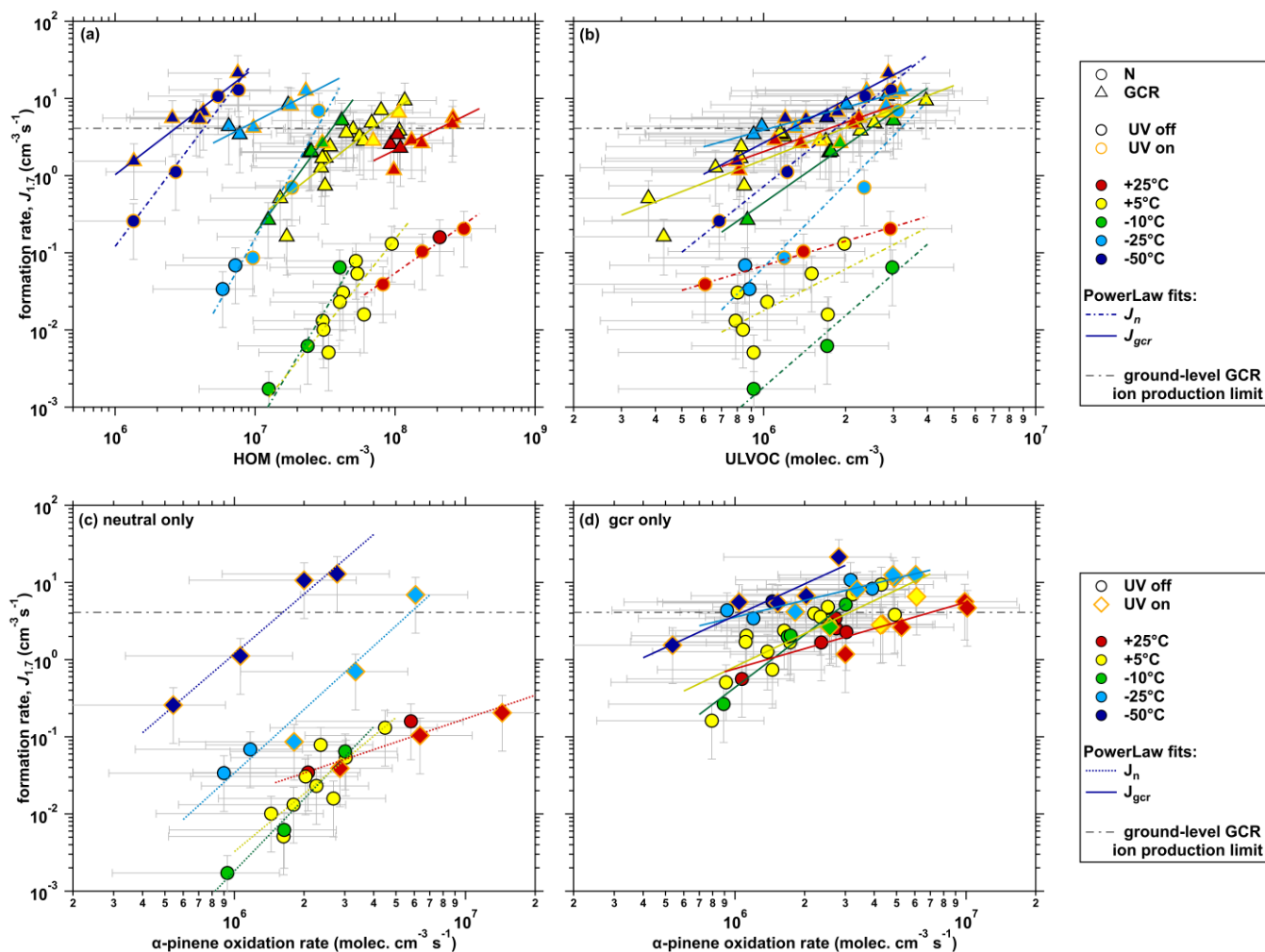
Figure 8: New-particle formation rate as function of HOM, ULVOC and α-pinene oxidation rates for temperatures from − 50 to + 25 °C.

Biogenic new-particle formation rates at 1.7nm ($J_{1.7}$) of α-pinene oxidation versus (a) total measured gas-phase HOMs and (b) ULVOC concentration under ion-free (J_n, circles) and charged (J_{gcr}, triangle) conditions. The color of the marker points specifies the investigated temperatures. The experimental conditions are 200 – 2000 pptᵥ of α-pinene, 37 – 48 ppbᵥ of O₃,

relative humidity 40 - 90 % and $< 1 \cdot 10^5$ cm^{-3} H$_2$SO$_4$. Black marker strokes indicate dark conditions, while orange marker strokes represent UV-illuminated conditions. Dash-dotted lines (J_n) and solid lines (J_{gcr}) are power law fits of the distinct ion condition for each of the individual temperatures. The dark gray dash-dotted line indicates the upper limit of ion-induced nucleation (J_{iin}) from the GCR ion-pair production rate at ground level. (c) Neutral (J_n) and (d) charged (J_{gcr}) new-particle formation rate as a function of α-pinene oxidation rates at dark condition (circles) and UV-illuminated condition (diamonds). The bars indicate 1σ total errors, although the overall systematic scale uncertainty of + 78 %/− 68 % on the HOM mixing ratios is not shown.