- 1 Dear Referees,
- 2 Thanks for giving us an opportunity to revise our manuscript (acp-2019-1045). We appreciate your positive and con-
- 3 structive comments. We have studied these comments carefully and make revisions on the manuscript. These com-
- 4 ments and the corresponding replies are listed below.
- 5 The referee's comments are highlighted by gray. The symbol ">>" quotes the original texts in the manuscript. Followed
- by the comments are our responses (normal texts) and current texts in the manuscript (leaded by **line number**). Some
- 7 important revisions are colored by red. The revised manuscript with track changes are attached at the end of this file.
- 8 With regards,

11

21

9 Shuqi Yan, Bin Zhu\*, and all co-authors.

# 12 **Replies to Referee#1**

- 13 **1.** Line 37-38: As indicated in IPCC AR5, "aerosol-cloud-radiation interactions" is suggested to be rephrased as "aero-
- sol-radiation interaction" and "aerosol-cloud interaction" separately.
- 37-38: ...which are called as aerosol-cloud-radiation interactions...
- 16 Thank you for this valuable suggestion. We have corrected it.

# 17 Line 38-39 (Introduction)

- 18 ...which are called as aerosol-radiation and aerosol-cloud interactions...
- 19 **Line 458 (Table 1)**
- 20 Aerosol-cloud-radiation interactions --> Aerosol-cloud and aerosol-radiation interactions
- 22 **2.** Line 41: "Many"->"Previous"
- >> Line 41: Many studies have analysed...

- 24 Thank you for this valuable suggestion. We have changed all the "many studies" to "previous studies" (Line 35, 42,
- 25 219).

- **3.** Line 45: "lower supersaturation"??
- 28 »Line 45: As a result, urban areas commonly experience higher temperatures and lower vapour contents. These condi-
- 29 tions induce a lower supersaturation that is unfavourable for fog formation.
- Thank you for this valuable suggestion. We have corrected this sentence.

### Line 46 (Introduction)

...These conditions induce a lower relative humidity supersaturation that is unfavourable for fog formation.

33

31

32.

- 4. Lines 64-65: Some important references are missing regarding the observational evidences of aerosol boomerang
- 35 effect in China, e.g., Wang et al., AE 2015, doi: 10.1016/j.atmosenv.2015.04.063; Guo et al., GRL 2017, doi:
- 36 10.1002/2017GL073533; Liu et al., Sci. Rep. 2019, doi:10.1038/s41598-019-44284-2.
- 37 »Lines 64-65: However, if aerosol concentration exceeds a certain threshold, this promoting effect disappears (Quan
- 38 et al., 2011) or even turns into a suppressing effect due to the strong vapour competition (Koren et al., 2008;
- 39 Rangognio, 2009).
- 40 Thank you for this valuable suggestion. We have added these references to the end of this sentence.

### Line 68 (Introduction)

- 42 .....or even turns into a suppressing effect due to the strong vapour competition (Guo et al., 2017; Koren et al.,
- 43 2008; Liu et al., 2019; Rangognio, 2009; Wang et al., 2015).

44

- 45 **5.** Lines 69-71: I notice that the work by Yan et al. JGR (2019) mentioned here is also from the same research group.
- Also, it occurs to me that the motivation seems a little confused: Since previous work has "quantitatively" proved...,

47 why the authors attempt again to "quantitatively" confirm by model simulation of a fog event. Two "quantitatively" is 48 redundant. Therefore, this sentence is suggested to be rephrased as follows: e.g. Our recent observational work (Yan et 49 al., 2019) indicated a decreasing trend in fog days, and ..." 50 »Lines 69-71: Yan et al. (2019) analysed decadal trends of fog days and quantitatively proved that the inhibiting ef-51 fects of urbanization outweigh the promoting effects of aerosols on fog during the mature urbanization stage. Their 52 study inspires us to quantitatively confirm the roles of urbanization and aerosols..... 53 Thank you for this valuable suggestion. The redundant "quantitatively" is deleted. We have corrected this sentence. 54 **Line 72-76 (Introduction)** 55 Our recent observational work (Yan et al., 2019) indicated a decreasing trend in fog days, and the inhibiting ef-56 fects of urbanization outweigh the promoting effects of aerosols on fog during the mature urbanization stage. This 57 study aims to quantitatively confirm the roles of urbanization and aerosols..... 58 59 **6.** Line 75: "facilitates"-> "is expected to facilitate" 60 >> Line 75: This work facilitates the understanding of..... 61 Thank you for this valuable suggestion. We have corrected these sentences. 62 **Line 79 (Introduction)** 63 This work is expected to facilitate the understanding of... Line 305 (Conclusions) 64 65 This study is expected to facilitate a better understanding of... 66 67 7. Line 85: Something is suggested to be mentioned concerning Section 4 immediately after "Section 3.6 discusses the

68

69

rationality and reliability of the results."

»Line 85: Section 3.6 discusses the rationality and reliability of the results.

70 Thank you for this valuable suggestion. We have added something after it. 71 Line 89-90 (Introduction) 72. Section 3.6 discusses the rationality and reliability of the results. Section 4 concludes the findings of this study. 73 74 8. Line 90: it is suggested to clarify which city you are referring to? Since the reader cannot easily get any info from 75 either text or Figure 1b. 76 »Line 90: SX is ... approximately 30 km away from the nearest large city (Fig. 1b). 77 Thank you for this valuable suggestion. We have clarified this city, Huainan. It has been marked in Figure 1b. 78 **Line 95 (Section 2.1)** 79 SX is ... approximately 30 km away from the nearest large city, Huainan (Fig. 1b). 80 81 **9.** Line 97: "replace"-> "used to replace" 82 »Line 97: The data are resampled from 500 m to 30 arc-seconds (approximately 1 km) and replace the geological data 83 of the WRF model. 84 Thank you for this valuable suggestion. We have corrected this sentence. 85 **Line 102 (Section 2.1)** 86 The data are resampled from 500 m to 30 arc-seconds (approximately 1 km) and used to replace the geological 87 data of the WRF model. 88 89 10. Lines 160-165: The logic seems a little problematic: since the fog holes are mainly caused by urbanization, as

demonstrated in the references in this paragraph (aerosol effect is not mentioned and is supposed to not be the focus

here), why you mentioned the effect of aerosol pollution. It is generally thought that urbanization effect tends to reduce

90

- 92 LWP whereas aerosol tends to accumulate the formation of fog. The combined effect is highly dependent on the com-
- 93 peting effect of the two factors. Here it is not accurate to argue that both of them "reducing the LWP or advancing the
- 94 dissipation of fog".
- 95 »Line 160-165: ..... We assume that urbanization and air pollution could have profound effects on fog by reducing
- the LWP or advancing the dissipation of fog.
- 97 Thank you for this valuable suggestion. We agree that fog holes are mainly caused by urbanization, not by aerosols.
- 98 We aimed to express that "the combined effects of urbanization and aerosols lead to fog holes", not "both of them lead
- 99 to fog holes". To avoid the problem you mentioned, we have corrected the last sentence.

# Line 169-170 (Section 3.1)

- We assume that <u>urbanization</u> could have profound effects on fog by reducing the LWP or advancing the dissipa-
- tion of fog, and the role of aerosols on fog is weaker than that of urbanization.
- 104 11. Section 3.2: What are the criteria for you to determine a fog event from model-simuated LWP, which is required to
- be clarified here.

100

103

108

114

- Thank you for this valuable suggestion. The criteria for fog is LWP>2 g/m<sup>2</sup> (Jia et al., 2018). We have clarified it in
- 107 Section 3.2.

### **Line 173 (Section 3.2)**

- Satellite cloud image and modelled LWP (>2 g/m²) can represent the observed and simulated fog zone (Jia et al.,
- 110 2018).

### 111 **References**

- Jia, X., Quan, J., Zheng, Z., Liu, X., Liu, Q., He, H., and Liu, Y.: Impacts of anthropogenic aerosols on fog in North China Plain, J. Ge-
- ophys. Res.-Atmos., 124, 252–265, https://doi.org/10.1029/2018jd029437, 2018.

**12.** Section 3.3: The authors attempted to discuss the complicated non-monotonic effect of aerosol on fog formation by differing the emission rate, which is not very common. Why not used the aerosol concentration or CCN or Na that can well represent the real atmospheric pollution level for the time period investigated here? I am curious of the actual CCN or Na concentration for the experiment of u0e3?

Thank you for this valuable suggestion. We agree that CCN can better represent the air pollution level. The  $CCN_{0.1}$  concentration of each experiment is marked in the new Figure 8, because the supersaturation in fog is commonly less than 0.1% (Mazoyer et al., 2016). The  $CCN_{0.1}$  of current pollution level (u0e3) is 570 cm<sup>-3</sup>.

# Line 510-514 (Figure 8)

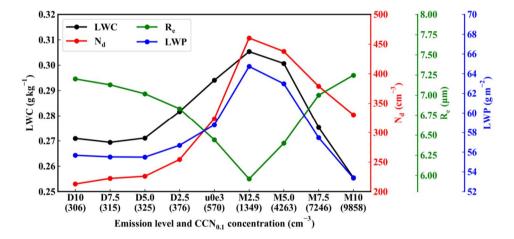


Figure 8. Relationships of the microphysical parameters (LWC,  $N_d$ ,  $R_e$  and LWP) with emission level and  $CCN_{0.1}$  concentrations. These parameters are the time-height averages (time average for the LWP) in fog.

### References

Mazoyer, M., Burnet, F., Roberts, G. C., Haeffelin, M., & Elias, T. (2016). Experimental study of the aerosol impact on fog microphysics. Atmospheric Chemistry and Physics, 1-35.

13. Lines 215-216: It will be misleading for the statement "the current pollution level in China is still located in the promoting regime rather than the suppressing regime of fog occurrence". Both ideal simulation (e.g., Rosenfeld et al. Science, 2008) or observational studies (Wang et al., AE 2015; Guo et al., GRL 2017) indicated that the tipping point

- tends to occur at AOD of 0.3-0.4 or CCN concentration of 1200/cm3. Recent observational work by Ilan Koren et al.
- (Science, 2014) suggested the cloud and precipitation is most sensitive to aerosol over the South Ocean. By compari-
- son, the average AOD from MODIS in East China is on average much larger than 0.6, irrespective of the meteorologi-
- cal conditions.
- 37 »Lines 215-216: The aerosol concentration of the transition point (experiment M2.5) is higher than that of u0e3 (Fig.
- 8), revealing that the current pollution level in China is still located in the promoting regime rather than the suppress-
- ing regime of fog occurrence.
- Thank you for this valuable suggestion. The  $CCN_{0.4}$  of u0e3 is 6023 cm<sup>-3</sup>, higher than  $CCN_{0.4}$ =1200 cm<sup>-3</sup> that revealed
- by Rosenfeld et al. (Science, 2008). We agree that the AOD value of East China is larger than 0.6. It seems that the
- current pollution level could suppress fog rather than promotes fog.
- The studies you listed mostly aim at convective clouds. Aerosols affect convective clouds through two competing
- mechanisms: 1) invigorating convection by promoting vapor condensation. 2) suppressing convection by blocking so-
- lar radiation and reducing surface heat flux. Under polluted conditions (AOD>0.3 or CCN<sub>0.4</sub>>1200 cm<sup>-3</sup>), the suppress-
- ing effect outweighs the invigoration effect, so the turning point occurs (Koren et al. Science, 2008; Rosenfeld et al.,
- Science, 2008). This suppressing effect does not exist in fog because fog commonly formed at night. Therefore, the
- turning point in fog might occur later than that in convective clouds. In North China Plain where air pollution is
- thought to be more serious, a case study by WRF-Chem also indicates that fog properties (e.g., LWC, N<sub>d</sub> and LWP)
- ineaght to be more sense, a case start of write enem also materials and log properties (e.g., 2 we, 1 and 2 will
- increase monotonically when emission intensity varies from 0.05-fold to 1-fold. It is consistent with our statement "the
- current pollution level in China is still located in the promoting regime rather than the suppressing regime of fog oc-
- 152 currence".

- 153 The above discussions have been included at the end of Section 3.4. Additionally, some statements are given in a more
- 154 cautious manner.

## Line 223-231 (Section 3.4)

- Rosenfeld et al. (2008) revealed that the turning point of boomerang pattern in convective clouds is  $CCN_{0.4} =$
- 157 1200 cm<sup>-3</sup>. The CCN<sub>0.4</sub> of u0e3 is 6023 cm<sup>-3</sup>, which seems to suppress fog. Aerosols affect convective clouds
- through two competing mechanisms: 1) invigorating convection by promoting vapour condensation. 2) suppress-
- ing convection by blocking solar radiation and reducing surface heat flux. Under polluted conditions (AOD>0.3

160	or CCN <sub>0.4</sub> >1200 cm <sup>-3</sup> ), the suppressing effect outweighs the invigoration effect, so the turning point occurs	
161	(Koren et al., 2008; Rosenfeld et al., 2008). This suppressing effect does not exist in fog because fog commonly	
162	formed at night. Therefore, the turning point in fog might occur later than that in convective clouds. In North Chi-	
163	na Plain where air pollution is thought to be more serious, a case study by WRF-Chem also indicates that fog	
164	properties (e.g., LWC, $N_d$ and LWP) increase monotonically when emission intensity varies from 0.05-fold to 1-	
165	fold.	
166	Line 220-222 (Section 3.4)	
167	revealing that possibly indicating that the current pollution level is still located in the promoting regime	
168	Line 25 (Abstract)	
169	the current pollution level in China is could be still below the critical aerosol concentration that suppresses fog.	
170		
171	<b>14.</b> Lines 205-206: times is redundant and should be removed.	
172	»Lines 205-206: For example, the name M2.5 means multiplying by 2.5 times; the name D10 means dividing by 10	
173	times.	
174	Thank you for this valuable suggestion. We have corrected this sentence.	
175	Line 210 (Section 3.4)	
176	For example, the name M2.5 means multiplying by 2.5; the name D10 means dividing by 10.	
177		
178	<b>15.</b> Line 448: it is better to indicate the year of 2017 following 03 January in the figure caption.	
179	»Line 448: Figure 2. The performance of the simulated fog zone at 08:00 03 January.	
180	Thank you for this valuable suggestion. We have changed two figure captions.	

The performance of the simulated fog zone at 08:00 03 January 2017....

Line 475 (Figure 2)

181

## Line 482 (Figure 3)

Two sub-regions (a and b) with obvious fog holes on the Himawari 8 image at 11:00 03 January 2017....

**16.** Figure 3: it is suggested to show the major cities in the regions of interest shown in panel a and b, given the convenience to better understand the fog hole induced by urban. Besides, two subregions in Figure 3 is better to be marked in Figure 1 or 2.

Thank you for this valuable suggestion. The subregions of interest are marked in the new Figure 2. The cities with fog holes are marked in the new Figure 3.

# Line 473-478 (Figure 2)

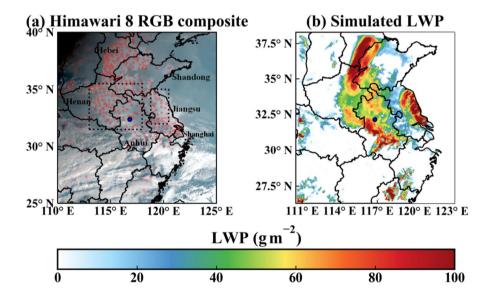


Figure 2. The performance of the simulated fog zone at 08:00~03 January 2017. (a) Himawari 8 RGB composite cloud image overlaid with the MICAPS observation sites (light red dots) at which fog was observed (relative humidity > 90 % and VIS < 1 km). (b) Simulated LWP distribution. Only LWC below 1500 m are integrated. The blue dots are the SX site. The two dashed rectangles in (a) are the subregions of interest in Fig. 3.

# Line 480-486 (Figure 3)

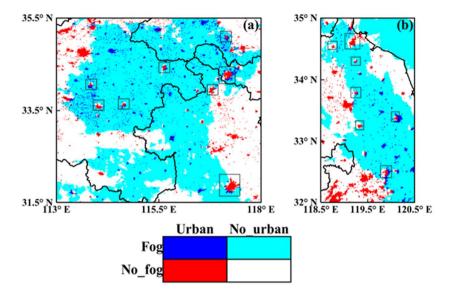


Figure 3. Two sub-regions (a and b) with obvious fog holes on the Himawari 8 image at 11:00~03 January 2017. The fog zone, which is represented by albedo > 0.45 (at  $0.64~\mu m$ ) and brightness temperature > 266~K (at  $12.4~\mu m$ ) (Di Vittorio et al., 2002), is marked with cold colours (blue or cyan). The urban areas are marked with blue or red. The red and white pixels surrounded or semi-surrounded by cold colours are fog holes, and among these pixels, the red pixels indicate the fog holes over urban areas. Some of the cities with fog holes are marked by rectangles.

# Replies to Referee#2

We appreciate your valuable suggestions of our manuscript.

208

209

206

207

# Replies to Referee#3

- 210 **1.** The focus of this study is on the radiative fog. However, there are different fog formation in atmosphere. For example, the advection fog formation is often occurred in the coast of eastern China. The Authors should highlight that under different fog conditions (i.e., radiative fog or advection fog, etc.) what is the effects of the urbanization and aerosol particles on the fog formation.
- Thank you for this valuable suggestion. We agree that radiation fog and advection fog are two major fog types in China. They can occur in both inland and coastal areas. Gu et al. (2019) revealed the occurring frequencies of different fog types in Shanghai, a coastal city, during the past three decades. The major fog type is radiation fog (38.3%), followed by advection fog (27.7%) and advection-radiation fog (23.4%). Therefore, we infer that the dominant fog type in in-
- 218 land areas and coastal areas is radiation fog, which should be attracted more attention.
- 219 Compared with radiation fog which usually occurs under stagnant weather conditions, advection fog is associated with
  - synoptic forcing, i.e., advection of a moist air mass with contrasting temperature properties with respect to the underly-
- 221 ing surface (Gultepe et al., 2007). The role of synoptic forcing should be considered when studying the effects of ur-
- banization and aerosols on advection fog, which is more complex than radiation fog. Zhong et al. (2017) indicated that
- 223 urbanization and aerosols have nonsignificant effects on convective precipitation when the synoptic forcing is strong.
- Therefore, this study focuses on radiation fog to study the effects of urbanization and aerosols.

### References

- Gu, Y., Kusaka, H., van Doan, Q., and Tan, J.: Impacts of urban expansion on fog types in Shanghai, China: Numerical experiments by WRF model, Atmos. Res., 220, 57–74, https://doi.org/10.1016/j.atmosres.2018.12.026, 2019.
- Zhong, S., Qian, Y., Zhao, C., Leung, R., Wang, H., Yang, B., Fan, J., Yan, H., Yang, X., and Liu, D.: Urbanization-induced urban heat island and aerosol effects on climate extremes in the Yangtze River Delta region of China, Atmos. Chem. Phys., 17, 5439–5457, https://doi.org/10.5194/acp-17-5439-2017, 2017.

220

232 2. Is this study suitable for the most of large cities in eastern China? 233 Thank you for this valuable suggestion. In the reply to comment, we infer that the dominant fog type in inland areas 234 and coastal areas of eastern China is radiation fog. Under the unified leadership of national government, most of the 235 cities in eastern China experience the similar development pattern, i.e., the expansion of urban areas is commonly ac-236 companied by increasing aerosol pollution. So we believe our results of radiation fog are suitable for most of the cities 237 in eastern China. 238 239 3. Some important references are missing. For example, Tie et al (2017) studied the important feedback of atmospheric 240 moister on the aerosol pollution in eastern China, which should state in the instruction. 241 References 242 Tie, X., R.J. Huang, J.J. Cao, O. Zhang, Y.F. Cheng, H. Su, D. Chang, U. Pöschl, T. Hoffmann, U. Dusek, G. H. Li, D. R. Worsnop, C. D. 243 O'Dowd, Severe Pollution in China Amplified by Atmospheric Moisture, Sci. Rep. 7: 15760 DOI:10.1038/s41598-017-15909-1, 2017. 244 Tie, XX, X. Long, GH Li, SY Zhao, JJ Cao, JM Xu, Ozone enhancement due to photo-dissociation of nitrous acid in eastern China, Atmos. 245 Chem. Phys., 19,11267-11278, 2019. 246 Thank you for this valuable suggestion. We have added the two references and some texts. 247 **Line 53-54 (Introduction)** 248 Aerosols exert sophisticated impacts on fog through direct (radiation) effects and indirect (microphysical) effects

(Khain and Pinsky, 2018). Aerosols attenuate shortwave radiation, influencing PBL structure and the vertical pro-

file of moisture and aerosols (Tie et al., 2017, 2019), which can alter the formation and dissipation condition of

249

250

251

fog....

# To what extents do urbanization and air pollution affect fog?

- Shuqi Yan<sup>1,2,3,4</sup>, Bin Zhu<sup>1,2,3,4,\*</sup>, Yong Huang<sup>5,6</sup>, Jun Zhu<sup>7</sup>, Hanqing Kang<sup>1,2,3,4</sup>, Chunsong Lu<sup>1,2,3,4</sup>, Tong Zhu<sup>8</sup>
- 3 Collaborative Innovation Centerre on Forecast and Evaluation of Meteorological Disasters, Nanjing University of Information Science & Technology, Nanjing, China
- 5 <sup>2</sup>Key Laboratory for Aerosol-Cloud-Precipitation of China Meteorological Administration, Nanjing University of Information
- 6 Science & Technology, Nanjing, China
- 7 3Key Laboratory of Meteorological Disaster, Ministry of Education (KLME), Nanjing University of Information Science &
- 8 Technology, Nanjing, China
- 9 <sup>4</sup>Special test field of National Integrated meteorological observation, Nanjing University of Information Science & Tech-
- 10 nology, Nanjing, China

1

- 11 Sanhui Meteorology Institute, Key Lab of Atmospheric Science and Remote Sensing Anhui Province, Hefei 230031, China
- 12 <sup>6</sup>Shouxian National Climatology Observatory, Shouxian 232200, China
- <sup>7</sup>Xiangshan Meteorological Bureau, Xiangshan 315700, China
- 14 8IMSG at NOAA/NESDIS/STAR, 5830 University Research Ct., College Park, MD 20740, USA
- 16 Correspondence to: Bin Zhu (binzhu@nuist.edu.cn)
- 17 Abstract. The remarkable development of China has resulted in rapid urbanization (urban heat island and dry island) and 18 severe air pollution (aerosol pollution). Previous studies demonstrate that these two factors have either suppressing or pro-19 moting effects on fog, but what are the extents of their individual and combined effects? In this study, a dense radiation fog 20 event in East China in January 2017 was reproduced by the WRF-Chem model, and the individual and combined effects of 21 urbanization and aerosols on fog (indicated by liquid water content (LWC)) are quantitatively revealed. Results show that 22 urbanization inhibits low-level fog, delays its formation and advances its dissipation due to higher temperatures and lower 23 saturations. In contrast, upper-level fog could be enhanced because of the updraft-induced vapour convergence. Aerosols 24 promote fog by increasing LWC, increasing droplet concentration and decreasing droplet effective radius. Further experi-25 ments show that the current pollution level in China is-could be still below the critical aerosol concentration that suppresses 26 fog. Urbanization influences fog to a larger extent than do aerosols. When urbanization and aerosol pollution are combined, 27 the much weaker aerosol promoting effect is counteracted by the stronger urbanization suppressing effect on fog. Budget 28 analysis of LWC reveals that urban development (urbanization and aerosols) alters LWC profile and fog structure mainly by 29 modulating condensation/evaporation process. Our results infer that urban fog will be further reduced if urbanization keeps 30 developing and air quality keeps deteriorating in the future.

### 1 Introduction

During the past five decades, China has achieved remarkable developments, accompanied by strong anthropogenic activities (rapid urbanization and severe air pollution). Urbanization and air pollution have significantly affected climate change, monsoons, air quality, fog, clouds and precipitation (e.g., Li et al., 2016; Li et al., 2017). Many Previous studies have linked the changes in clouds and precipitation to urbanization and aerosols. Urbanization destabilizes the boundary layer, which triggers strong updrafts and invigorates convection (e.g., Rozoff et al., 2003; Shepherd, 2005). Aerosols modify the macroscopic, microphysics, thermodynamics and radiative properties of clouds through complicated pathways, which are called as aerosol cloud radiation aerosol-radiation and aerosol-cloud interactions and have been systematically reviewed by Fan et al. (2016), Rosenfeld et al. (2014), Tao et al. (2012), etc. Fog can be viewed as a cloud (Leng et al., 2014) that occurs near the surface. Land use features and aerosol properties may instantly affect fog, so fog is more sensitive to anthropogenic activities than other types of clouds are (Zhu and Guo, 2016). Many Previous studies have analysed the effects of urbanization and aerosols on fog, mostly in segregated manners.

Urbanization is featured with urban heat island (UHI) and dry island (UDI) effects. The urban surface has a lower albedo, which reduces the reflected solar radiation and enhances heat storage. Urban expansion decreases the coverage of cropland, water bodies and forestland, which reduces the sources of water vapour. As a result, urban areas commonly experience higher temperatures and lower vapour contents. These conditions induce a lower relative humidity supersaturation that is unfavourable for fog formation (Gu et al., 2019). In the long-term scale, urban fog days are reported to decrease significantly (e.g., Guo et al., 2016; LaDochy, 2005; Sachweh and Koepke, 1995; Shi et al., 2008; Yan et al., 2019). Although UHI and UDI inhibit near-surface fog, the upward motions can promote upper-level fog (Li et al., 2011; Niu et al., 2010b). Surface roughness and thermal circulation cause strong updrafts (Rozoff et al., 2003), which transfer water vapour aloft and cause wet island phenomenon in the upper-level (Kang et al., 2014). The fog at that altitude may be subsequently enhanced.

Aerosols exert sophisticated impacts on fog through direct (radiation) effects and indirect (microphysical) effects (Khain and Pinsky, 2018). Aerosols attenuate shortwave radiation, influencing PBL structure and the vertical profile of moisture and aerosols (Tie et al., 2017, 2019), which can alter the formation and dissipation condition of fog. Scattering aerosols block downwelling solar radiation in the daytime, thus delaying the dissipation and elongating the duration of fog (Shi et al., 2008; Maalick et al., 2016). Although they increase downwelling longwave radiation at night, scattering aerosols have negligible effects on the fog formation time (Stolaki et al., 2015; Maalick et al., 2016). The role of absorbing aerosols like BC on fog depends on its residence height. If BC resides above the fog layer, BC causes a dome effect (Ding et al., 2016) which blocks solar radiation and prevents the dissipation of fog (Bott, 1991). If BC resides within the fog layer, BC heats fog droplets and accelerates the dissipation of fog (Maalick et al., 2016). The aerosol indirect effect on cloud is addressed as one of the most

批注 [yansq1]: Referee#1\_Comment2

批注 [vansq2]: Referee#1 Comment1

批注 [yansq3]: Referee#1\_Comment2

批注 [yansq4]: Referee#1\_Comment3

批注 [yansq5]: Referee#3\_Comment3

uncertain factors in the IPCC report (IPCC, 2013). Aerosol concentration has a two-fold effect on fog, which is called as the boomerang pattern (Koren et al., 2008). Under saturation conditions, increasing aerosols commonly result in more CCNs. It promotes activation and condensation, yielding more but smaller droplets and increasing cloud water content (Fan et al., 2018; Rosenfeld et al., 2008). These changes have two kinds of positive feedback on fog (Maalick et al., 2016): more droplets cause stronger radiative cooling at fog top and enhance condensation (Jia et al., 2018); smaller droplet size inhibits sedimentation and the depletion of cloud water (Zhang et al., 2014). However, if aerosol concentration exceeds a certain threshold, this promoting effect disappears (Quan et al., 2011) or even turns into a suppressing effect due to the strong vapour competition (Guo et al., 2017; Koren et al., 2008; Liu et al., 2019; Rangognio, 2009; Wang et al., 2015). Additionally, large-scale aerosol pollution can change weather patterns and affect large-scale fog formation conditions (Niu et al., 2010a). Ding et al. (2019) found that the dome effects of BC induce a land-sea thermal contrast and generate a cyclonic anomaly over coastal areas. This anomaly results in more vapor transported inland and strengthened advection-radiation fog.

批注 [yansq6]: Referee#1\_Comment4

Our recent observational work (Yan et al., 2019) indicated a decreasing trend in fog days, and Yan et al. (2019) analysed decadal trends of fog days and quantitatively proved that the inhibiting effects of urbanization outweigh the promoting effects of aerosols on fog during the mature urbanization stage. Their study inspires us This study aims to quantitatively comfirm confirm the roles of urbanization and aerosols in a dense fog event by an online-coupled synoptic and air quality model, WRF-Chem. This event is a radiation fog event with weak synoptic forcing (detailed in Sect. 3.1), so the effects of urbanization and aerosols should be obvious. Determining the quantitative extents of urbanization effect, aerosol effect and their combined effect is an interesting topic, which has barely been studied previously to the best of our knowledge. This work sexpected to facilitates the understanding of how anthropogenic activities affect the natural environment, fog (cloud) physics and aerosol-cloud interactions near the surface.

批注 [yansq7]: Referee#1\_Comment5

批注 [yansq8]: Referee#1\_Comment6

increasing aerosol pollution caused by urban expansion. Air pollution refers to aerosols and is indicated by anthropogenic emissions because aerosol concentration is highly proportional to emission intensity. Liquid water content (LWC) and cloud/fog droplet number concentration ( $N_d$ ) are two important parameters representing fog intensity and visibility. Following previous studies (e.g., Ding et al., 2019; Gu et al., 2019; Jia et al., 2018; Maalick et al., 2016; Yang et al., 2018), we use LWC as the indicator of fog to reveal different characteristics of fog in different experiments. This study is organized as follows. The data, model and methods are described in Sect. 2. Section 3.1 overviews the fog event and provides preliminary evidence of how urban development affects fog. Section 3.2 evaluates the model performance. Sections 3.3 to 3.5 analyse the urbanization, aerosol and combined effects on fog. Section 3.6 discusses the rationality and reliability of the results. Section 4 concludes the findings of this study.

In this study, urbanization mainly refers to UHI and UDI associated with land use change and human activities, excluding the

批注 [yansq9]: Referee#1\_Comment7

#### Data, model and methods 2

#### 2.1 Data 92

91

93

103

104

105

106

107 108

109

110

111

112

113

114

115

116

117

118

The first data are the hourly automatic weather station data from the Shouxian National Climate Observatory (SX; 32.4° N, 94 116.8° E, 23 m) that are used to evaluate the model performance. SX is a rural site surrounded by vast croplands and is ap-95 proximately 30 km away from the nearest large city, Huainan (Fig. 1b). The data include horizontal visibility, temperature, 96 relative humidity, wind direction and speed. The second data are the Himawari 8 satellite data that are used to represent fog 97 area (https://www.eorc.jaxa.jp/ptree/index.html). Fog area is mainly indicated by the albedo at three visible bands: red (band 98 3, 0.64 µm), green (band 2, 0.51 µm) and blue (band 1, 0.47 µm). The third data are the 3-hourly data from the Meteorologi-99 cal Information Comprehensive Analysis and Process System (MICAPS) (Li et al., 2010) that are also used to represent the 100 fog area. The fourth data are the land use data from the Moderate Resolution Imaging Spectroradiometer Land Cover Type Version 6 data (MCD12Q1; https://lpdaac.usgs.gov/products/mcd12q1v006) in the year of 2017, the same as the simulation 101 period. The data are resampled from 500 m to 30 arc-seconds (approximately 1 km) and used to replace the geological data 102

批注 [yansq10]: Referee#1 Comment8

批注 [vansq11]: Referee#1\_Comment9

### **Model configuration**

of the WRF model.

model that considers the sophisticated interactions among various dynamic, physical and chemical processes (Chapman et al., 2009; Fast et al., 2006). WRF or WRF-Chem has been successfully used in simulating fog events (Jia and Guo, 2012; Jia and Guo, 2015; Jia et al., 2018) and exploring aerosol-cloud interactions (Fan et al., 2018). Two nest domains are set up (Fig. 1). The d01 domain has a size of 217×223 grids and a resolution of 6 km, covering the entire fog area of this event (Fig. 2a). The d02 domain has a size of 115×121 grids and a resolution of 2 km, covering SX and the adjacent areas. The land use data are replaced by MCD12Q1 data, which represent the latest condition.

The model used in this study is the WRF-Chem (V3.9.1.1) model. It is an online-coupled mesoscale synoptic and air quality

Fog simulation is highly sensitive to vertical grids (Gultepe et al., 2007). A fine vertical resolution with a proper lowest model level can better resolve turbulences, thus yielding a reasonable fog structure (Yang et al., 2019). Here, 42 vertical levels are established with the first five n values of 1.000, 0.999, 0.998, 0.997, 0.996. There are 25 levels below the boundary

layer (approximately 1500 m), and the lowest model level is approximately 8 m.

Fog simulation is also sensitive to physical schemes (Gu et al., 2019). Through numerous experiments, radiation, microphysics and boundary schemes are found to significantly influence the model performance, and the boundary layer scheme plays a decisive role (Naira Chaouch et al., 2017). The radiation schemes are the RRTM longwave scheme and the Goddard

- shortwave scheme. The microphysical scheme is the Morrison double-moment scheme (Morrison et al., 2005). The boundary
- 120 layer scheme is the YSU 1.5-order closure non-local scheme, which yields better results than do any other schemes. The
- major schemes are listed in Tab. 1.
- 122 The model is driven by the highest resolution product (0.125°, approximately 13 km) of ECMWF data
- 123 (https://apps.ecmwf.int/datasets/data/interim-full-daily/levtype=sfc/). The anthropogenic emissions are derived from the
- 124 Multi-resolution Emission Inventory for China (MEIC) database (http://www.meicmodel.org). The simulation starts at
- 125 2017-01-01 08:00 and ends at 2017-01-03 14:00, with the first 24 hours as the spin-up period (all the times here are in local
- 126 time).

130

131

140

### 2.3 Sensitivity experiments

- 128 The study site is SX because only its visibility is observed hourly and is a multiple of 1 m, which is suitable for evaluating
- the model performance. To investigate the effects of urbanization and aerosols on fog, we change the land use and emission
  - intensity around SX. Four experiments, i.e., u0e0, u3e0, u0e3 and u3e3 are designed. The u0e0 is the base experiment, with
  - no urbanization and weak emission at SX. The u3e0 is set as the urbanization condition. The u0e3 is set as the polluted con-
- 132 dition. The u3e3 is set as the urban development condition (urbanization and pollution coexist). The experiment settings are
- listed in Tab. 2.
- 134 On the setting of urbanized condition, we replace the land use of SX as that of Hefei, the most urbanized city and the capital
- of Anhui Province. The downtown of Hefei has a built area of approximately 570 km<sup>2</sup>. Therefore, the 11x13 box centered on
- 136 SX (572 km²) is replaced by urban surface in the u3e0 and u3e3 experiments to represent the urbanization condition.
- 137 The downtown of Hefei has much higher emissions than SX. For example, the PM2.5 emission rate of Hefei is 40 times
- 138 higher than that of SX. To represent the polluted condition, the emission intensity of the aforementioned box is set to be
- equal to that of downtown Hefei in the u0e3 and u3e3 experiments.

### 2.4 Calculating visibility

- 141 The LWC is the proxy of fog as mentioned above. Since the LWC is not observed, and visibility (VIS) is related to LWC, the
- 142 VIS is used to assess the model performance. VIS is not diagnosed by the model and can be parameterized by the function of
- LWC, N<sub>d</sub> or droplet effective radius (R<sub>e</sub>). Equation 1 (Kunkel, 1983) and 2 (Gultepe et al, 2006) are two parameterization
- 144 methods.

VIS 
$$[m] = 27 LWC[g cm^{-3}]^{0.88}$$
 (1)

$$VIS[m] = 1002(LWC[g cm-3] \cdot N_a[cm-3])^{-0.6473}$$
(2)

- Another parameterization method is based on the Mie theory (Gultepe et al., 2017). VIS is inverse proportional to atmos-
- pheric extinction at visible wavelength. The extinction coefficient of cloud water ( $\beta_c$ ) is

$$\beta_c [\text{km}^{-1}] = \frac{3Q_{ext} \rho_a LWC}{4\rho_w R_e} \times 10^6$$
(3)

- where  $\rho_a$  ( $\rho_w$ ) is the air (water) density in kg m<sup>-3</sup>, LWC is in g kg<sup>-1</sup>,  $R_e$  is in  $\mu m$ , and  $Q_{ext}$  is the extinction efficiency, which is
- 148 assumed to be 2 for cloud droplets.
- The atmospheric extinction  $(\beta)$  is also largely contributed by aerosols  $(\beta_a)$  and other types of hydrometeors. The model diag-
- noses  $\beta_a$  at 550 nm. No other types of hydrometeors occur in this fog case, so we assume  $\beta = \beta_a + \beta_c$ . Then VIS is determined
- by the Koschmieder rule (Koschmieder, 1924): VIS[m]= $3.912/\beta$ [km<sup>-1</sup>]×1000.
- During fog period (Fig. 4 shaded zone), the three methods nearly yield the same results (figure not shown), so the last meth-
- od is used to calculate the simulated VIS.

# 3 Results and discussions

154

155

156

164

# 3.1 Overview of the fog event

### 3.1.1 Formation condition and lifetime

- 157 From 01 to 06 January 2017, East China is dominated by zonal circulation, with weak trough, ridge, pressure gradient and
- 158 atmospheric diffusion (Zhang and Ma, 2017). Under this stable weather pattern, the accumulation of pollutants and water
- 159 vapour promote the occurrence of fog-haze events. From the evening of 02 January to the noon of 03 January, a dense fog
- event occurs in wide regions of East China. The fog reaches its peak at 08:00 03 January, covering south Hebei, east Henan,
- 161 west Shandong, Anhui, Jiangsu and Shanghai (Fig. 2a). Figure 4a shows the temporal variation of visibility at SX. The fog
- 162 forms at 18:00 02 January and dissipates at 12:40 03 January. This is a radiation fog which is promoted by strong radiative
- 163 cooling at night and weak easterly water vapour transport from northwest Pacific (Zhu et al., 2019).

### 3.1.2 Preliminary evidence of urban development affecting fog

Lee (1987) and Sachweh and Koepke (1995) observed "fog holes" over urban areas on satellite images. Here, fog hole means

the low liquid water path (LWP) region within the fog region, which is visualized as pixels with weak fog (high visibility) or clear sky surrounded by dense fog. These holes demonstrate that urban development (urbanization and aerosols) has a clearing effect on fog. In this fog event, fog holes are also present over urban areas on the Himawari 8 image at 11:00 03 January (Fig. 3). We assume that urbanization and air pollution could have profound effects on fog by reducing the LWP or advancing the dissipation of fog, and the role of aerosols on fog is weaker than that of urbanization.

批注 [yansq12]: Referee#1\_Comment10

### 3.2 Model evaluation and simulations

The model performance is evaluated by comparing the fog spatial coverage. Satellite cloud image and modelled LWP 22 g m<sup>-2</sup> can represent the observed and simulated fog zone, respectively (Jia et al., 2018). Figure 2 shows the Himawari 8 visible cloud image and the simulated LWP distribution at 08:00. The light white pixels and light red dots indicate the observed fog area. The model well captures the fog in south Hebei, east Henan, west Shandong, Anhui, Jiangsu and Shanghai.

The model performance is also evaluated by comparing the visibility and other basic parameters at the SX site (Fig. 4). Seen from the visibility, the simulated fog forms at 19:30, 1.5 h later than the observation, and dissipates at 12:20, 30 min earlier than the observation. During the fog period, the simulated visibility agrees well with the observation. The other parameters such as temperature, wind speed and relative humidity are also effectively reproduced by the model, with relative small RMSEs of 0.8 K, 0.7 m/s and 5.9 %, respectively. Overall, the model well captures the spatial feature and temporal evolution of the fog.

### 3.3 Urbanization effects

From different sensitivity experiments (u3e0, u0e3 and u3e3), we can deduce the extents of the separate or combined effects of urbanization and aerosols on fog. Figure 5 compares the LWC between u0e0 and u3e0. The general results are: (1) Before 02:00, urbanization leads to a decreasing LWC in all layers. Fog forms on the surface at 22:30 in u3e0, 3 h later than in u0e0. (2) After 02:00, the LWC decreases in the low-level while it increases in the upper-level. Fog dissipates at 10:50 in u3e0, 1.5 h earlier than in u0e0. To better explain the LWC difference, its profiles are shown in Fig. 6. At 23:00, although fog has formed in u3e0, the fog is rather weak compared with u0e0, which is caused by the higher temperature (Fig. 6f) and lower saturation associated with UHI and UDI. At 02:00, fog develops in u3e0, but its intensity (the value of LWC) cannot reach the same level as that in u0e0.

An interesting phenomenon is the opposite change of LWC in the low-level and upper-level after 02:00. This phenomenon can be explained by the role of updrafts. The increasing roughness length and extra warming in urban conditions could trigger horizontal wind convergence (Fig. S1) and the enhanced updrafts (Fig. 5c). The stronger updrafts in u3e0 affect condensations of the condensation of the condensat

批注 [yansq13]: Refere#1\_Comment11

sation via two possible pathways: (1) the vertical transport of vapour  $(-w\frac{\partial q}{\partial z})$  and vertical convergence/divergence  $(-q\frac{\partial w}{\partial z})$  redistribute water vapour and affect condensation; (2) the adiabatic cooling promotes condensation. The role of the first pathway is measured by vertical vapour flux divergence  $(\frac{1}{g}\frac{\partial(qw)}{\partial z})$ . At 05:00, u3e0 shows a stronger vapour convergence above 110 m (Fig. 6h), and the LWC increases above 130 m (Fig. 6c). At 08:00, u3e0 shows a stronger vapour convergence above 130 m (Fig. 6i), and the LWC increases above 170 m (Fig. 6d). Therefore, it is possible that the adiabatic cooling and updraft-induced vapour flux convergence increase the vapour content and promote condensation in the upper-level, while the fog in the low-level is suppressed by the divergence of vapour flux. At 11:00, fog disappears at the ground in u3e0 likely due to the higher temperature (Fig. 6j). In summary, the UHI, UDI and updrafts alter the profile of LWC and reduce the LWP most of the time (Fig. 5c), and the decreasing LWP in the daytime can explain why fog holes occur above urban areas (Fig. 3).

### 3.4 Aerosol effects

Figure 7 compares the LWC between u0e0 and u0e3. The formation time, dissipation time of fog and fog top show almost no changes. The LWC increases at almost all layers in the polluted condition. Accordingly, the LWP also increases (Fig. 7c). It is probable that the current pollution level of China always promotes fog occurrence. To testify whether the u0e3 is below the transition point of the boomerang pattern, eight additional experiments (D10, D7.5, D5, D2.5, M2.5, M5, M7.5 and M10) are performed. These experiments are the same as u0e3, except that the emissions around SX (the black box in Fig. 1b) are multiplied (the "M" prefix) or divided (the "D" prefix). For example, the name M2.5 means multiplying by 2.5 times, the name D10 means dividing by 10 times.

Figure 8 compares the LWC, N<sub>d</sub>, R<sub>e</sub> and LWP among the nine emission-variant experiments. All the four parameters show the boomerang pattern, which demonstrates that the model is able to simulate the dual effects of aerosols. Below u0e3, the four parameters monotonically vary with emission level or CCN concentration, indicating that aerosol pollution could always promote fog. This phenomenon is because stronger emissions produce more aerosols and CCN. Under saturation conditions, the larger amount of CCN boost activation and yield a higher N<sub>d</sub>. The higher N<sub>d</sub> reduces R<sub>e</sub> and inhibits autoconversion and sedimentation (Twomey, 1977); thus, this situation decreases the depletion of fog water and increases the LWC. This promoting effect has been confirmed by many-previous model studies (e.g., Maalick et al., 2016; Stolaki et al., 2015) and observations (e.g., Chen et al., 2012; Goren and Rosenfeld, 2012). The aerosol CCN<sub>0.1</sub> concentration of u0e3 (570 cm<sup>-3</sup>) is lower than that of the turning point (experiment M2.5) (1349 cm<sup>-3</sup>) is higher than that of u0e3 (Fig. 8), revealing possibly indicating that the current pollution level in China (u0e3) is still located in the promoting regime rather than the suppressing regime of fog occurrence, which is also found by Jia et al. (2018).

批注 [yansq14]: Referee#1 Comment14

批注 [yansq15]: Referee#1 Comment2

Rosenfeld et al. (2008) revealed that the turning point of boomerang pattern in convective clouds is CCN<sub>0.4</sub> = 1200 cm<sup>-3</sup>. The CCN<sub>0.4</sub> of u0e3 is 6023 cm<sup>-3</sup>, which seems to suppress fog. Aerosols affect convective clouds through two competing mechanisms: 1) invigorating convection by promoting vapour condensation. 2) suppressing convection by blocking solar radiation and reducing surface heat flux. Under polluted conditions (AOD>0.3 or CCN<sub>0.4</sub>>1200 cm<sup>-3</sup>), the suppressing effect outweighs the invigoration effect, so the turning point occurs (Koren et al., 2008; Rosenfeld et al., 2008). This suppressing effect does not exist in fog because fog commonly formed at night. Therefore, the turning point in fog might occur later than that in convective clouds in North China Plain where air pollution is thought to be more serious; a case study by WRF-Chem also indicates that fog properties (e.g., LWC, N<sub>d</sub> and LWP) increase monotonically when emission intensity varies from 0.05-fold to 1-fold.

批注 [yansq16]: Refere#1\_Comment13

### 3.5 Combined effects of urbanization and aerosols

- Figure 9 compares the LWC between u0e0 and u3e3. The u3e3-induced change is quite similar to but not the same as the u3e0-induced change. The time-height average of absolute change of LWC induced by u3e0, u0e3 and u3e3 are 0.120, 0.019, 0.124 g kg<sup>-1</sup>, respectively. This result indicates that urbanization affects fog to a larger extent than do aerosols; when urbanization and aerosols are combined, the effect of aerosols is indiscernible. The LWP is also significantly suppressed in the day-
- time, and the promoting effect of aerosols in Fig. 7c is indiscernible in Fig. 9c. To further explain the changes in LWC, we perform budget analysis of the LWC to determine which physical processes are the dominant contributors.
- 239 In WRF, the budget of LWC is composed of the following items,

$$\frac{\partial q_c}{\partial t} = -\underbrace{\left(u\frac{\partial}{\partial x} + v\frac{\partial}{\partial y} + w\frac{\partial}{\partial z}\right)q_c}_{\text{adv}} + \left(\frac{\partial q_c}{\partial t}\right)_{\text{PBL}} + \left(\frac{\partial q_c}{\partial t}\right)_{\text{micro}} + \left(\frac{\partial q_c}{\partial t}\right)_{\text{cumu}}$$
(4)

- where  $q_c$  is LWC, and the subscripts denote advection, boundary layer, microphysical and cumulus processes, respectively.
- The microphysical tendency is further decomposed into the following items,

$$\left(\frac{\partial q_c}{\partial t}\right)_{\text{micro}} = \left(\frac{\partial q_c}{\partial t}\right)_{\text{cold}} + \left(\frac{\partial q_c}{\partial t}\right)_{\text{auto}} + \left(\frac{\partial q_c}{\partial t}\right)_{\text{accr}} + \left(\frac{\partial q_c}{\partial t}\right)_{\text{sedi}} + \left(\frac{\partial q_c}{\partial t}\right)_{\text{cond/evap}} \tag{5}$$

- where the subscripts denote cold phase processes, autoconversion, accretion, sedimentation and condensation/evaporation,
- 243 respectively.

223

224 225

226 227

228

229

230

231

232

244 All the processes regarding precipitation and cold phase (the cumu, cold, auto and accr subscripts) are not analysed because

no precipitation occurs, and the temperature is above  $0^{\circ}C$  in the simulated fog (figure not shown). The sum of microphysical (condensation/evaporation and sedimentation), boundary layer and advection tendencies is equal to the LWC distribution, so the contributions of other physical processes can be safely ignored.

We can also infer that to what extents the various physical processes affect fog through the sensitivity experiments (u3e0, u0e3 and u3e3). Additional aerosols weakly influence these processes (Fig. S2 right column) and subsequently result in weak LWC change (Fig. 7c). Compared with aerosols, urbanization effect is much more considerable (Fig. S3 right column); it dominantly accounts for the variation in physical tendencies from u0e0 to u3e3 (Fig. 10 right column). In u3e3 condition, urban development (urbanization and aerosols) induces different magnitude of changes in different physical tendencies. The relative magnitudes are 52.1, 38.3 and 9.6 % for the microphysical, boundary layer and advection processes, respectively, indicating that microphysics is most susceptible to urban development and contributes most to the LWC change. Among various microphysical processes, condensation/evaporation contributes most (72.7 %) to the change in microphysical tendency (Fig. 11 right column). The above results indicate that urban development affects the LWC mainly by modulating the condensation/evaporation process. Since u3e3 condition still witnesses higher temperatures and stronger updrafts (figure not shown), the notable variation in condensation/evaporation tendency induced by u3e3 can also be attributed to the predomi-

### 3.6 Discussions

As mentioned above, urbanization influences fog to a larger extent than do aerosols; the LWC in fog does not vary substantially with pollution level. This section discusses the rationality and reliability of our results through mechanism analysis and observational evidence.

nant role of UHI, UDI and updrafts. The mechanism has been analysed in Sect. 3.3.

- The sensitivity of cloud properties to aerosols depends on aerosol concentration and saturation environment. In convective clouds with intense upward motions and high saturations, the response of cloud properties to additional aerosols is significant ("aerosol-limited regime") (Fan et al., 2018). However, in fog with much weaker updrafts and lower saturations, this response could be more sensitive to vapour content rather than aerosol concentration ("vapour-limited regime"). It possibly implies that the LWC in fog varies slightly with pollution level but considerably with saturation condition that related to urbanization. Our results reveal that the time-height average LWC varies within the extent of 0.07g kg<sup>-1</sup> when emission intensity varies within two orders of magnitude (Fig. 8). This relative weak response of the LWC to pollution level is also reported by Jia et al. (2018).
  - In terms of observational evidence, Yan et al. (2019) revealed that fog days in polluted regions of East China have decreased since the 1990s. Through quantitative analysis, the promoting effects of aerosols are weakening, while the suppressing ef-

fects of urbanization are enhancing and dominantly cause this decrease. Sachweh and Koepke (1995) also claimed that the hindering effects of urbanization outweigh the promoting effects of aerosols on fog in southern Germany. Additionally, satellite images present discernible fog holes above urban areas (Fig. 3) (Lee, 1987; Sachweh and Koepke, 1995). Therefore, these observational evidence support the model results that the promoting effect of aerosols is counteracted by the hindering effect of urbanization. We believe that the results can also be applied to other <u>large</u> cities in China because these cities commonly witness strong UHI, UDI and severe air pollution.

## 4 Conclusions

- A dense radiation fog event occurred in East China from 02 to 03 January 2017. Satellite images show that fog holes occur over urban areas, demonstrating the remarkable effects of urbanization and air pollution on fog. Hence, the mechanism is investigated by the WRF-Chem model. The model well captures the spatial coverage and temporal evolution of the fog. Furthermore, the separate and combined effects of urbanization (refers to UHI and UDI) and air pollution (refers to aerosols) on fog (indicated by the LWC) are revealed, and the extents of these effects are quantitatively determined. Results show that:
  - Urbanization redistributes the LWC profile by the UHI, UDI effect and updrafts. The updrafts may be caused by surface roughness and extra warming. The UHI and UDI suppress low-level fog, delay its formation by 3 h, and advance its dissipation by 1.5 h. However, the upper-level fog could be enhanced due to the updraft-induced adiabatic cooling and vapour flux convergence. Urbanization reduces the LWP most of the time, and this reduction in the daytime can explain why fog holes are present above urban areas on satellite images.
  - Aerosols promote fog mainly by changing microphysical properties. The increasing emissions (aerosol concentration) produce more CCN and fog droplets, which decreases R<sub>e</sub> and inhibits sedimentation, thus leading to a higher LWC. Further sensitivity experiments show that the current pollution level in China is-could be still below the transition point of the boomerang pattern that suppresses fog. The macroscopic properties such as fog top and lifetime remain nearly unchanged.
  - The role of urbanization far overweighs that of aerosols. Therefore, when they act together, the urbanization effect is dominant, and the aerosol effect is indiscernible. Budget analysis of LWC shows that increasing aerosols influence various physical processes to a lesser extent, while urbanization influences these processes to a larger extent, eventually leading to a substantial LWC change in urban development condition (urbanization and aerosols). In this condition, the comparisons among various physical processes reveal that microphysics dominates the change in LWC, and condensation/evaporation dominates the change in microphysical tendency. This result highlights the importance of condensation/evaporation process in modulating the LWC profile and fog structure.

Mechanism analysis and the observational evidence support our key finding that urbanization influences fog to a much larger extent than do aerosol pollution. Therefore, we believe our results are reasonable and robust in radiation fog events without strong synoptic forcings, and the results can also be applied to other large cities in China due to the similar urban development patterns. This study is expected to facilitates a better understanding of how anthropogenic activities affect the natural environment, fog (cloud) physics and aerosol-cloud interactions near the surface. We can also infer the future change of fog occurrence. Under the traditional urban development pattern, i.e., urbanization keeps developing and air quality keeps deteriorating, urban fog occurrence will be further reduced. Code and data availability. Some of the data repositories have been listed in Sect. 2. The other data, model outputs and codes can be accessed by contacting Bin Zhu via binzhu@nuist.edu.cn. Author contributions. SY performed the model simulation, data analysis and manuscript writing. BZ proposed the idea, supervised this work and revised the manuscript. YH provided the observation data at the SX site. JZ processed the observation data. HK offered helps to the model simulation. CL and TZ also contributed to the manuscript revision. Competing interests. The authors declare that they have no conflict of interest. Acknowledgments. We are grateful to the High Performance Computing Center of Nanjing University of Information Science and Technology for doing the numerical calculations in this work on its blade cluster system. We thank American Journal Experts (AJE) for the English language editing. Financial support. This work is supported by the National Key Research and Development Program (2016YFA0602003) and the National Natural Science Foundation of China (91544229, 41575148, 41605091). References

302

303

304

305

306

307

308

309 310

311

312313

314

315

316 317

318 319

320

321

322 323

324

325 326

327

批注 [yansq17]: Refere#1\_Comment6

Abdul-Razzak, H. and Ghan, S. J.: A parameterization of aerosol activation 3. Sectional representation, J. Geophys. Res., 107,

AAC-1-AAC 1-6, https://doi.org/10.1029/2001jd000483, 2002.

- Bott, A.: On the influence of the physico-chemical properties of aerosols on the life cycle of radiation fogs, J. Aerosol. Sci., 21, 1–31, https://doi.org/10.1007/BF00119960, 1991.
- Chapman, E. G., Gustafson, W. I., Easter, R. C., Barnard, J. C., Ghan, S. J., and Pekour, M. S.: Coupling aerosol-cloud-radiative processes in the WRF-Chem model: Investigating the radiative impact of elevated point sources, Atmos. Chem. Phys., 9, 945–964, https://doi.org/10.5194/acp-9-945-2009, 2009.
- Chen, Y. C., Christensen, M. W., Xue, L., Sorooshian, A., Stephens, G. L., Rasmussen, R. M., and Seinfeld, J. H.: Occurrence of lower cloud albedo in ship tracks, Atmos. Chem. Phys., 12, 8223–8235, https://doi.org/10.5194/acp-12-8223-2012, 2012.
- Di Vittorio, A. V. and Emery, W. J.: An automated, dynamic threshold cloud-masking algorithm for daytime AVHRR images over land, IEEE Trans. Geosci. Remote Sensing, 40, 1682-1694, https://doi.org/10.1109/TGRS.2002.802455, 2002.
- Ding, A. J., Huang, X., Nie, W., Sun, J. N., Kerminen, V. M., Petäjä, T., Su, H., Cheng, Y. F., Yang, X. Q., Wang, M. H., Chi, X. G.,
   Wang, J. P., Virkkula, A., Guo, W. D., Yuan, J., Wang, S. Y., Zhang, R. J., Wu, Y. F., Song, Y., Zhu, T., Zilitinkevich, S., Kulmala, M.,
   and Fu, C. B.: Enhanced haze pollution by black carbon in megacities in China, Geophys. Res. Lett., 43, 2873–2879,
   https://doi.org/10.1002/2016g1067745, 2016.

342

343 344

345

346

347

348

349

350

351

352

355

356

357

358

359

360

- Ding, Q., Sun, J., Huang, X., Ding, A., Zou, J., Yang, X., and Fu, C.: Impacts of black carbon on the formation of advection-radiation fog during a haze pollution episode in eastern China, Atmos. Chem. Phys., 19, 7759–7774, <a href="https://doi.org/10.5194/acp-19-7759-2019">https://doi.org/10.5194/acp-19-7759-2019</a>, 2019.
- Fan, J., Rosenfeld, D., Zhang, Y., Giangrande, S. E., Li, Z., and Machado, L. A. T.: Substantial convection and precipitation enhancements by ultrafine aerosol particles, Science, 359, 411–418, https://doi.org/10.1126/science.aan8461, 2018.
- Fan, J., Wang, Y., Rosenfeld, D., and Liu, X.: Review of Aerosol–Cloud Interactions: Mechanisms, Significance, and Challenges, J. Atmos. Sci., 73, 4221–4252, <a href="https://doi.org/10.1175/JAS-D-16-0037.1">https://doi.org/10.1175/JAS-D-16-0037.1</a>, 2016.
- Fast, J. D., Gustafson, W. I., Easter, R. C., Zaveri, R. A., Barnard, J. C., Chapman, E. G., Grell, G. A., and Peckham, S. E.: Evolution of ozone, particulates, and aerosol direct radiative forcing in the vicinity of Houston using a fully coupled meteorology-chemistry-aerosol model, J. Geophys. Res., 111, https://doi.org/10.1029/2005jd006721, 2006.
- Goren, T. and Rosenfeld, D.: Satellite observations of ship emission induced transitions from broken to closed cell marine stratocumulus over large areas, J. Geophys. Res.-Atmos., 117, -, <a href="https://doi.org/10.1029/2012JD017981">https://doi.org/10.1029/2012JD017981</a>, 2012.
- Gu, Y., Kusaka, H., van Doan, Q., and Tan, J.: Impacts of urban expansion on fog types in Shanghai, China: Numerical experiments by WRF model, Atmos. Res., 220, 57–74, https://doi.org/10.1016/j.atmosres.2018.12.026, 2019.
  - Gultepe, I., Tardif, R., Michaelides, S. C., Cermak, J., Bott, A., Bendix, J., Müller, M. D., Pagowski, M., Hansen, B., Ellrod, G., Jacobs, W., Toth, G., and Cober, S. G.: Fog Research: A Review of Past Achievements and Future Perspectives, Pure Appl. Geophys., 164, 1121–1159, https://doi.org/10.1007/s00024-007-0211-x, 2007.
  - Gultepe, I., Müller, M. D., and Boybeyi, Z.: A New Visibility Parameterization for Warm-Fog Applications in Numerical Weather Prediction Models, J. Appl. Meteorol. Climatol., 45, 1469–1480, https://doi.org/10.1175/jam2423.1, 2006.
    - Gultepe, I., Milbrandt, J. A., and Zhou, B.: Marine fog: A review on microphysics and visibility prediction, in: Koračin D., Dorman C. (eds) Marine Fog: Challenges and Advancements in Observations, Modeling, and Forecasting, Springer, Cham, 50 pp., 2017.
- Guo, J., Su, T., Li, Z., Miao, Y., Li, J., Liu, H., Xu, H., Cribb, M., and Zhai, P.: Declining frequency of summertime local-scale precipitation over eastern China from 1970 to 2010 and its potential link to aerosols, Geophys. Res. Lett., 44, 5700–5708, https://doi.org/10.1002/2017GL073533, 2017.
- Guo, T., Zhu, B., Kang, Z., Gui, H., and Kang, H.: Spatial and temporal distribution characteristic of fog days and haze days from
   1960-2012 and impact factors over the Yangtze River Delta Region, China Environmental Science, 36, 961 969,
   https://doi.org/10.3969/j.issn.1000-6923.2016.04.001, 2016. [in Chinese]
- IPCC: Climate change 2013: The physical science basis, Contribution of Working Group I to the Fifth Assessment Report of the Intergov ernmental Panel on Climate Change, Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 1585 pp.,
   2013.

- Jia, X. and Guo X.: Impacts of Anthropogenic Atmospheric Pollutant on Formation and Development of a Winter Heavy Fog Event, Chinese Journal of Atmospheric Sciences, 36, 995–1008, https://doi.org/10.3878/j.issn.1006-9895.2012.11200, 2012. [in Chinese]
- Jia, X. and Guo, X.: Impacts of Secondary Aerosols on a Persistent Fog Event in Northern China, Atmospheric and Oceanic Science Letters, 5, 401–407, https://doi.org/10.1080/16742834.2012.11447022, 2015.
- Jia, X., Quan, J., Zheng, Z., Liu, X., Liu, Q., He, H., and Liu, Y.: Impacts of anthropogenic aerosols on fog in North China Plain, J. Geophys. Res.-Atmos., 124, 252–265, https://doi.org/10.1029/2018jd029437, 2018.
  - Kang, H., Zhu, B., Zhu, T., Sun, J., and Ou, J.: Impact of Megacity Shanghai on the Urban Heat-Island Effects over the Downstream City Kunshan, Bound.-Layer Meteor., 152, 411–426, <a href="https://doi.org/10.1007/s10546-014-9927-1">https://doi.org/10.1007/s10546-014-9927-1</a>, 2014.
  - Khain, A. P. and Pinsky, M.: Modeling: A Powerful Tool for Cloud Investigation, in: Physical processes in clouds and cloud modeling, Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 98 pp., 2018.
  - Koren, I., Martins, J. V., Remer, L. A., and Afargan, H.: Smoke invigoration versus inhibition of clouds over the Amazon, Science, 321, 946–949, https://doi.org/10.1126/science.1159185, 2008.
- 383 Koschmieder, H.: Therie der horizontalen sichtweite, Beitr Phys.d.freien Atm, 12, 171–181, 1924.

378

379

380

381

382

386

387

390

391

392

395

396

401

402

406

407

- Kunkel, B. A.: Parameterization of Droplet Terminal Velocity and Extinction Coefficient in Fog Models, J. Appl. Meteorol., 23, 34–41, https://doi.org/10.1175/1520-0450(1984)023<0034:PODTVA>2.0.CO;2, 1983
  - LaDochy, S.: The Disappearance of Dense Fog in Los Angeles: Another Urban Impact?, Phys. Geogr., 26, 177–191, https://doi.org/10.2747/0272-3646.26.3.177, 2005.
- 388 Lee, T. F.: Urban clear islands in California central valley fog, Mon. Weather Rev., 115, 1794–1796, 389 https://doi.org/10.1175/1520-0493(1987)1152.0.CO;2, 1987.
  - Leng, C., Zhang, Q., Zhang, D., Xu, C., Cheng, T., Zhang, R., Tao, J., Chen, J., Zha, S., and Zhang, Y.: Variations of cloud condensation nuclei (CCN) and aerosol activity during fog-haze episode: a case study from Shanghai, Atmos. Chem. Phys., 14, 12499–12512, https://doi.org/10.5194/acp-14-12499-2014, 2014.
- Li, Y., Cao, L., Gao, S., and Luo, B.: The Current Stage and Development of MICAPS, Meteorological Monthly, 36, 50-55, 2010. [in Chinese]
  - Li, Z., Guo, J., Ding, A., Liao, H., Liu, J., Sun, Y., Wang, T., Xue, H., Zhang, H., and Zhu, B.: Aerosol and boundary-layer interactions and impact on air quality, Natl. Sci. Rev., 4, 810–833, https://doi.org/10.1093/nsr/nwx117, 2017.
- Li, Z., Lau, W. K. M., Ramanathan, V., Wu, G., Ding, Y., Manoj, M. G., Liu, J., Qian, Y., Li, J., Zhou, T., Fan, J., Rosenfeld, D., Ming, Y.,
  Wang, Y., Huang, J., Wang, B., Xu, X., Lee, S. S., Cribb, M., Zhang, F., Yang, X., Zhao, C., Takemura, T., Wang, K., Xia, X., Yin, Y.,
  Zhang, H., Guo, J., Zhai, P. M., Sugimoto, N., Babu, S. S., and Brasseur, G. P.: Aerosol and monsoon climate interactions over Asia,
  Rev. Geophys., 54, 866–929, <a href="https://doi.org/10.1002/2015RG000500">https://doi.org/10.1002/2015RG000500</a>, 2016.
  - Li, Z., Yang, J., Shi, C., and Pu, M.: Urbanization Effects on Fog in China: Field Research and Modeling, Pure Appl. Geophys., 169, 927–939, https://doi.org/10.1007/s00024-011-0356-5, 2011.
- Liu, H., Guo, J., Koren, I., Altaratz, O., Dagan, G., Wang, Y., Jiang, J. H., Zhai, P., and Yung, Y. L.: Non-Monotonic Aerosol Effect on precipitation in Convective Clouds over tropical oceans. Sci. Rep., 9, 1-7, https://doi.org/10.1038/s41598-019-44284-2, 2019.
   Maalick, Z., Kühn, T., Korhonen, H., Kokkola, H., Laaksonen, A., and Romakkaniemi, S.: Effect of aerosol concentration and absorbing
  - Maalick, Z., Kühn, T., Korhonen, H., Kokkola, H., Laaksonen, A., and Romakkaniemi, S.: Effect of aerosol concentration and absorbing aerosol on the radiation fog life cycle, Atmos. Environ., 133, 26–33, <a href="https://doi.org/10.1016/j.atmosenv.2016.03.018">https://doi.org/10.1016/j.atmosenv.2016.03.018</a>, 2016.
  - Morrison, H., Curry, J. A., and Khvorostyanov, V. I.: A new double-moment microphysics parameterization for application in cloud and climate models. Part I: Description, J. Atmos. Sci., 62, 1665–1677, <a href="https://doi.org/10.1175/JAS3446.1">https://doi.org/10.1175/JAS3446.1</a>, 2005.
- Naira Chaouch, Marouane Temimi, Michael Weston, and Hosni Ghedira: Sensitivity of the meteorological model WRF-ARW to planetary
  boundary layer schemes during fog conditions in a coastal arid region, Atmos. Res., 187, 106–127,
  https://doi.org/10.1016/j.atmosres.2016.12.009, available at: http://www.sciencedirect.com/science/article/pii/S0169809516307116,
  2017.
- Niu, F., Li, Z., Li, C., Lee, K., and Wang, M.: Increase of wintertime fog in China: Potential impacts of weakening of the Eastern Asian
   monsoon circulation and increasing aerosol loading, J. Geophys. Res., 115, <a href="https://doi.org/10.1029/2009jd013484">https://doi.org/10.1029/2009jd013484</a>, 2010a.

- 415 Niu, S., Lu, C., Yu, H., Zhao, L., and Lü, J.: Fog research in China: An overview, Adv. Atmos. Sci., 27, 639–662, 416 https://doi.org/10.1007/s00376-009-8174-8, 2010b.
- Rangognio, J.: Influence of aerosols on the formation and development of radiation fog, Atmos. Chem. Phys., 9, 17963–18019, https://doi.org/10.5194/acpd-9-17963-2009, 2009.
- Rosenfeld, D., Meinrat O. Andreae, Asmi, A., Chin, M., and Johannes Quaas: Global observations of aerosol-cloud-precipitation-climate interactions, Rev. Geophys., 52, 750–808, <a href="https://doi.org/10.1002/2013RG000441">https://doi.org/10.1002/2013RG000441</a>, 2014.

422

423

424

425

426

429

430

431

432

433

434

435

440

441

442

- Rosenfeld, D., Lohmann, U., Raga, G. B., O'Dowd, C. D., Kulmala, M., Fuzzi, S., Reissell, A., and Andreae, M. O.: Flood or drought: how do aerosols affect precipitation?, Science, 321, 1309–1313, https://doi.org/10.1126/science.1160606, 2008.
- Rozoff, C. M., Cotton, W. R., and Adegoke, J. O.: Simulation of St. Louis, Missouri, Land Use Impacts on Thunderstorms, J. Appl. Meteorol., 42, 716–738, https://doi.org/10.1175/1520-0450(2003)042<0716:SOSLML>2.0.CO;2, 2003.
- Sachweh, M. and Koepke, P.: Radiation fog and urban climate, Geophys. Res. Lett., 22, 1073–1076, <a href="https://doi.org/10.1029/95gl00907">https://doi.org/10.1029/95gl00907</a>, 1995.
- Shepherd, J. M.: A Review of Current Investigations of Urban-Induced Rainfall and Recommendations for the Future, Earth Interact., 9, 1-27, https://doi.org/10.1175/ei156.1, 2005.
  - Shi, C., Roth, M., Zhang, H., and Li, Z.: Impacts of urbanization on long-term fog variation in Anhui Province, China, Atmos. Environ., 42, 8484–8492, <a href="https://doi.org/10.1016/j.atmosenv.2008.08.002">https://doi.org/10.1016/j.atmosenv.2008.08.002</a>, 2008.
  - Stolaki, S., Haeffelin, M., Lac, C., Dupont, J. C., Elias, T., and Masson, V.: Influence of aerosols on the life cycle of a radiation fog event. A numerical and observational study, Atmos. Res., 151, 146–161, https://doi.org/10.1016/j.atmosres.2014.04.013, 2015.
  - Tao, W. K., Chen, J. P., Li, Z., Wang, C., and Zhang, C.: Impact of aerosols on convective clouds and precipitation, Rev. Geophys., 50, 6837, https://doi.org/10.1029/2011RG000369, 2012.
  - Tie, X., Huang, R., Cao, J., Zhang, Q., Cheng, Y., Su, H., Chang, D., Pöschl, U., Hoffmann, T., Dusek, U., Li, G., Worsnop, D., and
- 436 O'Dowd, C.: Severe Pollution in China Amplified by Atmospheric Moisture, Sci. Rep. 7, 15760, 437 <a href="https://doi.org/10.1038/s41598-017-15909-1">https://doi.org/10.1038/s41598-017-15909-1</a>, 2017.
- Tie, X., Long, X., Li, G., Zhao, S., Cao, J., and Xu, J.: Ozone enhancement due to photo-dissociation of nitrous acid in eastern China, Atmos. Chem. Phys., 19, 11267–11278, https://doi.org/10.5194/acp-19-11267-2019, 2019.
  - Twomey, S. A.: The Influence of Pollution on the Shortwave Albedo of Clouds, J. Atmos. Sci., 34, 1149–1154, https://doi.org/10.1175/1520-0469(1977)034<1149:tiopot>2.0.co;2, 1977.
  - Wang, F., Guo, J., Zhang, J., Huang, J., Min, M., Chen, T., Liu, H., Deng, M., and Li, X.: Multi-sensor quantification of aerosol-induced variability in warm clouds over eastern China, Atmos. Environ., 113, 1-9, <a href="https://doi.org/10.1016/j.atmosenv.2015.04.063">https://doi.org/10.1016/j.atmosenv.2015.04.063</a>, 2015
- Yan, S., Zhu, B., and Kang, H.: Long-term fog variation and its impact factors over polluted regions of East China, J. Geophys.
   Res.-Atmos., 124, 1741–1754, <a href="https://doi.org/10.1029/2018JD029389">https://doi.org/10.1029/2018JD029389</a>, 2019.
- 446 Yang, Y., Hu, X., Gao, S., and Wang, Y.: Sensitivity of WRF simulations with the YSU PBL scheme to the lowest model level height for a sea fog event over the Yellow Sea, Atmos. Res., 215, 253–267, https://doi.org/10.1016/j.atmosres.2018.09.004, 2019.
- Zhang, N. and Ma, X.: Analysis of the June 2018 Atmospheric Circulation and Weather, Meteorological Monthly, 43, 508–512, https://doi.org/10.7519/j.issn.1000-0526.2017.04.014, 2017. [in Chinese]
- Zhang, X., Musson-Genon, L., Dupont, E., Milliez, M., and Carissimo, B.: On the Influence of a Simple Microphysics Parametrization on
   Radiation Fog Modelling: A Case Study During ParisFog, Bound.-Layer Meteor., 151, 293–315,
   https://doi.org/10.1007/s10546-013-9894-y, 2014.
- 453 Zhu, B. and Guo, T.: Review of the Impact of Air Pollution on Fog, Advances in Meteorological Science and Technology, 6, 56–63, https://doi.org/10.3969/j.issn.2095-1973.2016.02.006, 2016. [in Chinese]
- 455 Zhu, J., Zhu, B., Huang, Y., An, J., and Xu, J.: PM2.5 vertical variation during a fog episode in a rural area of the Yangtze River Delta, 456 China, Sci. Total. Environ., 685, 555–563, https://doi.org/10.1016/j.scitotenv.2019.05.319, 2019.

# Table 1. Summary of major parameterization schemes.

Scheme	Option
Boundary layer	YSU
Longwave radiation	RRTM
Shortwave radiation	New Goddard
Microphysics	Morrison
Surface layer	MM5 similarity
Land surface	Noah
Urban surface	Urban canopy model
Gas phase chemistry	CBMZ
Aerosol chemistry	MOSAIC (4-bin)
Aerosol-cloud and aerosol-radiation interactions	All turned on
Aerosol activation	Abdul-Razzak and Ghan (2002)

批注 [yansq18]: Referee#1\_Comment1

Table 2. Settings of sensitive experiments. "N" represents no changes.

Case name	Description	Underlying surface	Anthropogenic emission
u0e0	base condition	N	N
u3e0	urbanization condition	the 11x13 grid centered on SX is replaced by urban surface	N
u0e3	polluted condition	N	the 11x13 grid centered on SX is replaced by the emission of Hefei downtown
u3e3	urbanization and polluted condition	same as u3e0	same as u0e3
Effect		Description	
u3e0-u0e0		urbanization effect	
u0e3-u0e0		aerosol effect	
u3e3-u0e0		urbanization and aerosol effect	

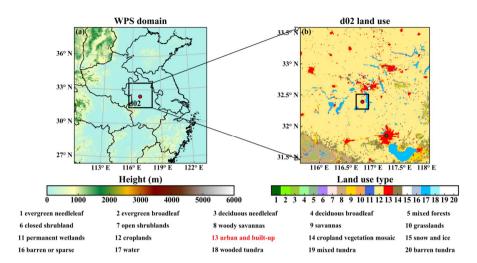


Figure 1. (a) The WRF domain overlaid with terrain height. (b) The land use distribution of domain d02. The green dot is Hefei, the capital of Anhui Province. The white dot is Huainan. The two red dots are the SX site. The land use and emissions of the  $22 \text{ km} \times 26 \text{ km}$  black box in the center of (b) will be altered in the sensitivity experiments.

批注 [yansq19]: Refere#1\_Comment8

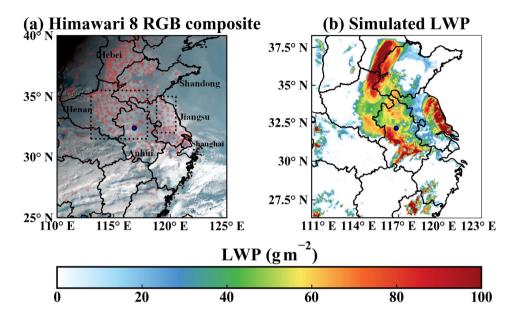


Figure 2. The performance of the simulated fog zone at 08:00 03 January 2017. (a) Himawari 8 RGB composite cloud image overlaid with the MICAPS observation sites (light red dots) at which fog was observed (relative humidity > 90 % and VIS < 1 km). (b) Simulated LWP distribution. Only LWC below 1500 m are integrated. The blue dots are the SX site. The two dashed rectangles in (a) are the subregions of interest in Fig. 3.

批注 [yansq20]: Referee#1\_Comment15

批注 [yansq21]: Referee#1\_Comment16

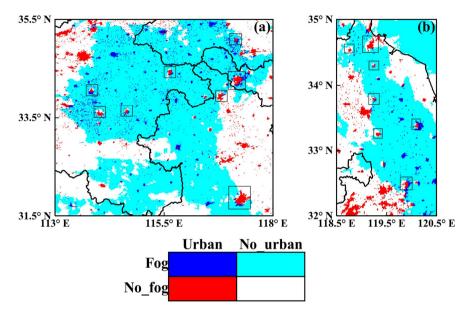


Figure 3. Two sub-regions (a and b) with obvious fog holes on the Himawari 8 image at 11:00~03~January~2017. The fog zone, which is represented by albedo > 0.45 (at  $0.64~\mu m$ ) and brightness temperature > 266~K (at  $12.4~\mu m$ ) (Di Vittorio et al., 2002), is marked with cold colours (blue or cyan). The urban areas are marked with blue or red. The red and white pixels surrounded or semi-surrounded by cold colours are fog holes, and among these pixels, the red pixels indicate the fog holes over urban areas. Some of the cities with fog holes are marked by rectangles.

批注 [yansq22]: Referee#1\_Comment16

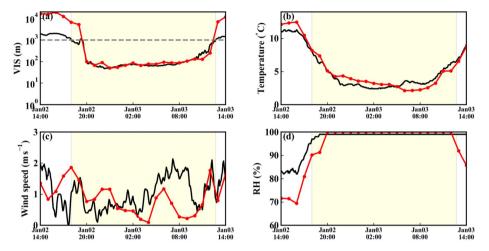


Figure 4. The performance of the simulated meteorological parameters at the SX site. (a) VIS. (b) air temperature. (c) 10-minute average wind speed. (d) Relative humidity (RH). The red dotted lines represent the model results, and the black lines are the observations. The fog period (VIS < 1 km and RH > 90 %) is shaded with light yellow.

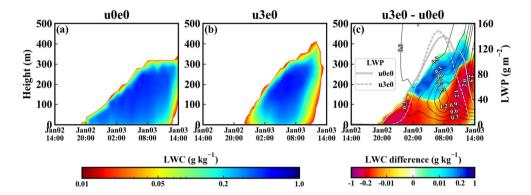


Figure 5. Time-height distribution of the LWC (g kg<sup>-1</sup>) in (a) u0e0 and (b) u3e0, and (c) is the urbanization effect (u3e0 minus u0e0) on LWC. The two white curves in (c) are the LWP. The black contour lines in (c) are the difference of vertical velocity (cm s<sup>-1</sup>) (u3e0 minus u0e0). Only the lines after 00:00 are shown for clarity.

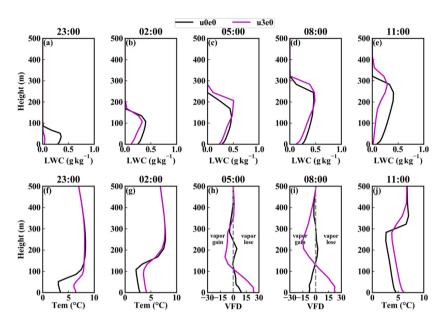


Figure 6. Profiles of the LWC (first row), temperature (Tem) (f, g, j) and vertical vapour flux divergence (VFD) (h, i)  $(g h^{-1} m^{-2} \cdot hpa^{-1})$  in u0e0 and u3e0 at different times.

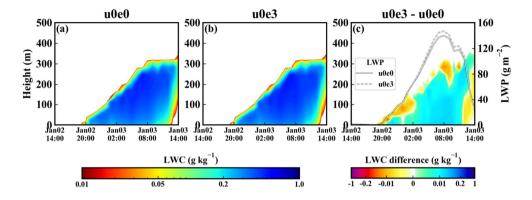


Figure 7. Similar to Fig. 5, but for the aerosol effect (u0e3 minus u0e0).

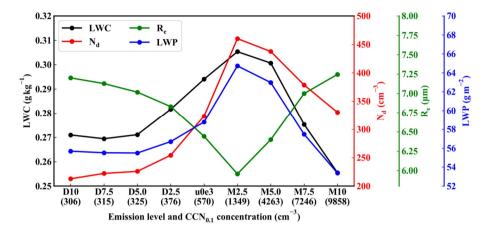
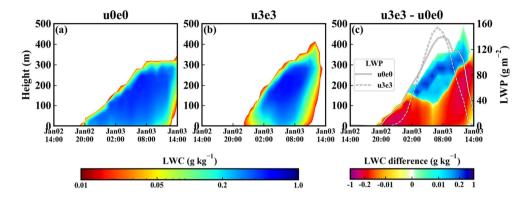


Figure 8, Relationships of the microphysical parameters (LWC, N<sub>d</sub>, R<sub>e</sub> and LWP) with emission level and CCN<sub>0.1</sub> concentrations. These parameters are the time-height averages (time average for the LWP) in fog, taking only non zero values into consideration.

批注 [yansq23]:  $CCN_{0.1}$  is marked under the corresponding experiments

批注 [yansq24]: Referee#1\_Comment12 & 13



 $Figure\ 9.\ Similar\ to\ Fig.\ 5,\ but\ for\ the\ combined\ effect\ of\ urbanization\ and\ aerosols\ (u3e3\ minus\ u0e0).$ 

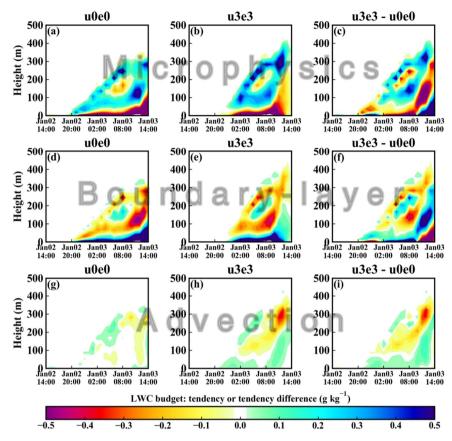


Figure 10. The combined effect of urbanization and aerosols (u3e3 minus u0e0) on various items of the LWC budget. The three rows are the  $\frac{1-\text{hour}}{\text{accumulated}}$  tendencies (g kg<sup>-1</sup>) of the microphysical, boundary layer, and advection processes.

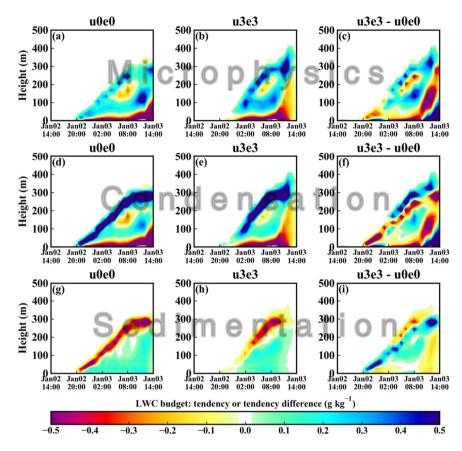


Figure 11. The combined effect of urbanization and aerosols (u3e3 minus u0e0) on various items of the microphysical tendency. The three rows are the  $\frac{1-\text{hour accumulated}\underline{\text{hourly}}}{\text{tendencies}}$  tendencies (g kg<sup>-1</sup>) of the microphysical, condensation/evaporation, and sedimentation processes.