

Our point-by-point responses are provided below. The referees' comments are italicized.

### **Response to Referee #1**

*Referee: This study aims to explore the reasons behind the elevated levels of ammonia observed over the Indo-Gangetic Plain. This is an important and scientifically relevant question, particularly since the ammonia burden has significant implications on inorganic aerosol concentrations over the region. The authors use the WRF-Chem model to investigate the physics and thermodynamics underlying the atmospheric fate of NH<sub>3</sub> and the resulting analysis provides useful insights into some of the factors driving the high concentrations over the region. For instance, the demonstrated sensitivity to increasing NO<sub>x</sub> and SO<sub>2</sub> emissions is an interesting result, particularly when contrasted to the case over the North China Plain.*

*While the dynamic physical transport and meteorology simulated by the model is validated and constrained, the approach chosen for the thermodynamic analysis largely hinges on an offline test of emissions and temperature sensitivities using the ISORROPIA model. The analysis would benefit significantly from a more rigorous on-line treatment of the thermodynamics that investigates the specific factors controlling NH<sub>3</sub> partitioning in greater detail. In addition, observational constraints (such as satellite ammonia columns) could be further leveraged to get a more quantitative estimate of model performance prior to interpreting the model output. With this in mind, I provide the following comments below and recommend that these issues be addressed prior to publication in ACP.*

**Response:** We would like to thank the referrer for your detailed and constructive comments. Please see our point-by-point reply below.

#### ***Referee: 1. Model Details, Validation and Uncertainties***

*In order to appropriately interpret the WRF-Chem analysis the authors should provide a more detailed discussion on the specifics of the ammonia simulation in WRF-Chem. For instance, is it possible to run ISORROPIA in an online configuration (partitioning at each time-step and explicitly simulating the aerosol species)? Unless I am misinterpreting the methods section, it appears that ISORROPIA is used only in an offline context. Given the spatial and temporal heterogeneity in the various factors that drive aerosol partitioning, running ISORROPIA in an online configuration would more appropriately explore the scientific questions outlined in this study.*

**Response:** In fact, the MOSAIC (Model for Simulating Aerosol Interactions and Chemistry) aerosol module embedded in WRF-Chem is used for online calculation of aerosol partitioning. Since WRF-Chem is a coupled model, variables such as air temperature cannot be arbitrarily perturbed. Thus, ISORROPIA-II is used offline to study the effect of air temperature on the gas-particle partitioning of NH<sub>3</sub>. As you suggested, we added more details about the WRF-Chem simulation in Sect 2.1.

**Revision:** (Page 2, Line 57-63) “CBM-Z (Carbon Bond Mechanism version Z) chemical mechanism (Zaveri and Peters, 1999) and MOSAIC (Model for Simulating Aerosol Interactions and Chemistry) aerosol module (Zaveri et al., 2008) were used for modeling photochemical reactions and aerosol processes (nucleation, growth and thermodynamic equilibrium), respectively. Wet deposition in the model includes both in-cloud and below-cloud scavenging. The below-cloud scavenging of aerosols by impaction and trace gases by mass transfer was based on the methods of Easter et al. (2004). Dry deposition was treated following the methods of Wesely (Wesely, 1989) for trace gases, and Binkowski and Shankar (Binkowski and Shankar, 1995) for aerosols.”

*Line 70: The authors have validated the simulated meteorology (wind speed, temperature, etc.) but do not validate the ammonia simulation itself. While I recognize this is challenging, it could potentially be done using satellite measurements (with the appropriate application of an averaging kernel) or surface measurements where available. Even a general estimate of how well the model captures ammonia variability and magnitude over the region would provide important context. Ideally, the different (non-transport) factors that dictate ammonia concentrations (namely – emissions, wet deposition, dry deposition and aerosol partitioning) would be constrained using observational data whenever available. In absence of such data, an explanation of the uncertainties associated with these various processes (and the steps that need to be taken to constrain them) is required in order to appropriately interpret the results.*

**Response:** It is inappropriate to quantitatively compare IASI NH<sub>3</sub> columns with modeled NH<sub>3</sub> columns. Because the IASI-NH<sub>3</sub> retrieval does not produce an averaging kernel to properly weight the model values, and using a typical averaging kernel will cause more uncertainty as there is a large day-to-day variability in the averaging kernels (Clarisse et al., 2009). The published observation data of NH<sub>3</sub> and NH<sub>4</sub><sup>+</sup> around 2010 over the IGP are almost all seasonal averages, and no monthly (or above) resolution data has been found. Since the simulated time range (June-August) is inconsistent with the time range of the seasonal average (April-June or March-June is summer; July-September or July-October is the monsoon season), the simulated value cannot be compared with the observation quantitatively. Nevertheless, as you suggested, we added a general evaluation of model simulation from NH<sub>3</sub> magnitude and spatial variations using IASI-NH<sub>3</sub> as a reference.

For the uncertainty brought by the dry and wet deposition, it is difficult to explore the uncertainty caused by this factor due to the lack of relevant observation data. However, the relevant researches have confirmed the relative accuracy of the existing dry and wet deposition parameterization in WRF-Chem (Easter et al., 2004; Wesely, 1989). In addition, for the impact of thermodynamic partitioning on ammonia concentrations, we have

studied the effect of temperature and humidity on aerosol partitioning which could further affect ammonia concentrations, see reply 5 and reply 7 for detail.

**Revision:** (Page 5, Line 146-152) “It is noted that the IASI NH<sub>3</sub> columns cannot be quantitatively compared to modeled NH<sub>3</sub> columns as the IASI NH<sub>3</sub> products do not provide information on the vertical sensitivity (averaging kernels) to properly weight the model values. Nonetheless, the simulated regional mean NH<sub>3</sub> total column over the IGP from the base case of  $8.8 \times 10^{16}$  molecules cm<sup>-2</sup> is close to the satellite-derived value ( $7.6 \times 10^{16}$  molecules cm<sup>-2</sup>), indicating that the model could generally capture the magnitude of NH<sub>3</sub> columns. Additionally, a broadly similar pattern was found in the NH<sub>3</sub> columns in the base run as in the satellite observations, both of which showed that NH<sub>3</sub> columns decrease along the IGP from northwest to southeast with the highest values in northwestern IGP (Figs. 1 and 2a)”

*Line 53: The simulations described here are spun-up over a 7-day period. A more detailed discussion about the initial concentrations assumed for the most important chemical species and their estimated lifetimes would provide useful context on whether the week-long period provides sufficient time to allow the longer-lived gas-phase species to equilibrate prior to the main simulation period.*

**Response:** To explore the effect of spin-up period on the simulation results, we extended the spin-up period from one week to one month. Results show that extending the spin-up period caused NH<sub>3</sub> columns over the IGP to change by only 0.007%, indicating that the 7-day spin-up would not affect the NH<sub>3</sub> simulation.

### **Referee: 2. Seasonal Analysis**

*Line 52: The authors classify June to August as the summer period. However, in India, this season is characterized by the monsoons (usually beginning in mid-June) which are associated with drastic changes to regional meteorology. This perhaps provides context for the statement on Line 198, given that high levels of precipitation and humidity are expected during the monsoon season. The Indian summer is usually thought to be between the months of April – May.*

**Response:** Accepted. As you suggested, we have corrected “summer” to “pre-monsoon and monsoon seasons”. The title and the relevant part of main text have been revised.

### **Referee: 3. Impact of Transport and Meteorology**

*Section 3.2 discusses the importance of various meteorological drivers (such as RH*

and temperature). However, given its importance in determining ammonia burdens, a more detailed discussion of the specific mechanisms that dictate aerosol partitioning under different meteorological conditions and the associated uncertainties in our understanding of these processes would add to the broader utility of this study.

**Response:** Accepted. To analyze the specific factors that affect  $\text{NH}_3$  concentrations under different meteorological conditions, we divided the IGP from northwest to southeast into three regions (namely western IGP, central IGP, and eastern IGP), and divided the study period into two seasons (pre-monsoon and monsoon). Analyses of the sensitivity of  $\text{NH}_3$  to temperature under different meteorological conditions have been added in Sect 3.2.

**Revision:** (Page 4, Line 98-101) “Sensitivity tests was firstly conducted based on the average of the entire IGP from June to August. Then, the IGP was divided equally from northwest to southeast into three regions (namely western IGP, central IGP, and eastern IGP), and the study period was divided into the pre-monsoon season (June) and the monsoon season (July to August). Sensitivity tests were conducted for the three regions under the two seasons.”

(Page 6, Line 182-186) “As shown in Figure 3c, the sensitivity of  $\text{NH}_3$  to temperature varies in different seasons and regions. Temporally, the sensitivity of  $\text{NH}_3$  to temperature during the monsoon season is generally higher than that during the pre-monsoon season. Spatially, the sensitivity of  $\text{NH}_3$  to temperature is highest over the eastern IGP, followed by the central IGP and the western IGP. The difference in the sensitivity of the  $\text{NH}_3$  to temperature may be caused by the difference of the initial  $\epsilon(\text{NH}_4^+)$  and temperature.”

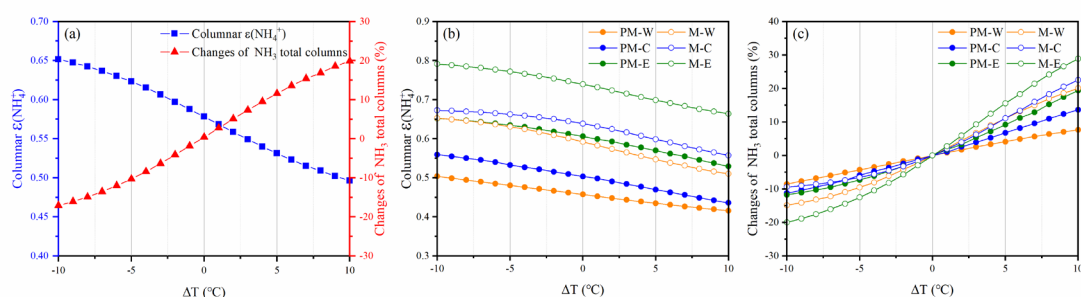


Figure 3. Columnar  $\epsilon(\text{NH}_4^+)$  and changes of  $\text{NH}_3$  total columns with the changes of temperatures predicted by ISORROPIA-II. (a) Mean columnar  $\epsilon(\text{NH}_4^+)$  and changes of  $\text{NH}_3$  total columns over the IGP from June to August 2010. (b) Columnar  $\epsilon(\text{NH}_4^+)$  and (c) changes of  $\text{NH}_3$  total columns over the western IGP during Pre-monsoon (PM-W), the central IGP during Pre-monsoon (PM-C), the eastern IGP during pre-monsoon (PM-E), the western IGP during monsoon (M-W), the central IGP during monsoon (M-C), the eastern IGP during monsoon (M-E).

**Referee:** 4. Use of the ISORROPIA module to access the impact of emissions and temperature

*Line 77: A more detailed overview of the ISORROPIA module would provide important context for the resulting analysis. If only applied in an offline context, it is possible that the analysis is not capturing various important (and non-linear) effects due to the spatial heterogeneity in gas phase and particle concentrations (along with the associated depositional losses at every timestep). While the authors provide an observational constraint (Line 142) to validate this approach, the differences in the model and observed partitioning ratio are significant (on the order of 30%). A more thorough comparison with the observational data would greatly benefit the analysis and serve as validation for some of the later conclusions. Additionally, the comparison of the regional mean to observational data over Delhi may not be appropriate, particularly given that NO<sub>x</sub> and SO<sub>2</sub> concentrations are likely much higher over the city.*

**Response:** In WRF-Chem, factors that affect NH<sub>3</sub> concentrations (including transport, dry deposition, wet deposition and gas-particle partitioning, etc.) are all calculated online. The MOSAIC (Model for Simulating Aerosol Interactions and Chemistry) aerosol module embedded in WRF-Chem is used for online calculation of aerosol partitioning.

As you suggested, we calculated the simulated  $\epsilon(\text{NH}_4^+)$  (partitioning ratios of NH<sub>4</sub><sup>+</sup> to total ammonia (TA, TA = NH<sub>3</sub> + NH<sub>4</sub><sup>+</sup>)) during monsoon over Delhi and compared it with observational data. The simulated  $\epsilon(\text{NH}_4^+)$  during monsoon over Delhi was 0.38, which is comparable to the observed  $\epsilon(\text{NH}_4^+)$  (0.39 in the 2011 monsoon season in Delhi (Singh and Kulshrestha, 2012)).

**Revision:** (Page 6, Line 163-164) “The simulated  $\epsilon(\text{NH}_4^+)$  in the 2010 monsoon in Delhi was 0.38, which is close to the observed  $\epsilon(\text{NH}_4^+)$  (0.39 in the 2011 monsoon season in Delhi) (Singh and Kulshrestha, 2012).”

*Line 155: The temperature sensitivity is an interesting result, particularly when contrasted to the SO<sub>2</sub>/NO<sub>x</sub> sensitivity. However, the approach here considers only a simplified case over the entire region. If the partitioning was conducted online (at every timestep), would it be reasonable to expect a different sensitivity to changes in temperature / SO<sub>2</sub> / NO<sub>x</sub>? A more detailed discussion about the non-linear, spatially dependent factors driving aerosol concentrations (and the heterogeneity in emissions, loss processes, thermodynamics, etc.) would provide more context with which to interpret these results.*

**Response:** The MOSAIC (Model for Simulating Aerosol Interactions and Chemistry) aerosol module embedded in WRF-Chem is used for online calculation of aerosol partitioning. Thus, the sensitivity of NH<sub>3</sub> to changes in SO<sub>2</sub>/NO<sub>x</sub> is the result of an online calculation.

Since WRF-Chem is a coupled model, variables such as air temperature cannot be arbitrarily perturbed. Thus, ISORROPIA-II is used offline to study the effect of air temperature on the gas-particle partitioning of NH<sub>3</sub>.

Given the temporal and spatial heterogeneity, we divided the IGP from northwest to southeast into three areas (namely western IGP, central IGP, and eastern IGP), and divided the study period into two seasons (pre-monsoon and monsoon). The sensitivity of  $\text{NH}_3$  to changes of temperature was analyzed for the three regions under the two seasons, see reply 5 for detail.

**Referee: 5. A discussion of other drivers of aerosol formation (particularly in the context of the monsoons)**

*Singh and Kulshrestha (2012), cited in this study, hypothesized that humidity during the monsoon season had a significant impact on  $\text{NH}_3$  partitioning. Could the authors discuss this in the context of their analysis? Given that the aim of this study is to establish the most salient drivers of the high  $\text{NH}_3$  concentrations (particularly during the selected monsoon period), a more detailed discussion of what determines the relative dominance of the different production, loss and partitioning mechanisms under various atmospheric conditions would provide important context with which to interpret the results of this analysis.*

**Response:** Accepted. Singh and Kulshrestha (2012) found that the percent fraction of  $\text{NH}_4^+$  over Delhi during monsoon was noticeably higher than that during the pre-monsoon and post-monsoon. Here, the model simulation also shows similar results. The simulated surface  $\epsilon(\text{NH}_4^+)$  over the western, central and eastern part of the IGP were 0.11, 0.13 and 0.24 during pre-monsoon and 0.26, 0.26 0.37 during monsoon. It is not difficult to find that the surface  $\epsilon(\text{NH}_4^+)$  during the monsoon season is significantly higher than that during the pre-monsoon season, and the surface  $\epsilon(\text{NH}_4^+)$  generally increases from northwest to southeast along the IGP. The spatiotemporal variations of  $\epsilon(\text{NH}_4^+)$  are consistent with the spatiotemporal variations of RH, indicating that RH is an important factor affecting the  $\text{NH}_3$  partitioning.

However, it is difficult to distinguish the contribution of various factors such as emission, chemical conversion and transport. As far as we understand, the high  $\text{NH}_3$  loading over the IGP appears to be the joint result of high  $\text{NH}_3$  emissions, weak chemical loss, and weak horizontal diffusion. Intensive agricultural activities resulted in relatively high  $\text{NH}_3$  emissions over the IGP. The low  $\text{SO}_2$  and  $\text{NO}_x$  emissions and high temperature limited the gas-to-particle partitioning of  $\text{NH}_3$ . In addition, orographic and meteorological conditions is conducive to the accumulation of  $\text{NH}_3$  over the IGP.

**Revision:** (Page 8, Line 225-231) “The simulated surface  $\epsilon(\text{NH}_4^+)$  over the western, central and eastern part of the IGP were 0.11, 0.13 and 0.24 during pre-monsoon and 0.26, 0.26 0.37 during monsoon. It is not difficult to find that the surface  $\epsilon(\text{NH}_4^+)$  during the monsoon season is significantly higher than that during the pre-monsoon season, and the surface  $\epsilon(\text{NH}_4^+)$  generally

increases from northwest to southeast along the IGP. Besides, the columnar  $\epsilon(\text{NH}_4^+)$  shows similar spatiotemporal variations with the surface  $\epsilon(\text{NH}_4^+)$  (Figure 3b). The spatiotemporal variations of  $\epsilon(\text{NH}_4^+)$  are consistent with the spatiotemporal variations of RH (Figure S5a), indicating that RH is an important factor affecting the  $\text{NH}_3$  partitioning.”

## **References**

- Binkowski, F. S., and Shankar, U.: The Regional Particulate Matter Model: 1. Model description and preliminary results, *J. Geophys. Res. Atmos.*, 100, 26191-26209, <https://doi.org/10.1029/95JD02093>, 1995.
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- Singh, S., and Kulshrestha, U. C.: Abundance and distribution of gaseous ammonia and particulate ammonium at Delhi, India, *Biogeosciences*, 9, 5023-5029, <https://doi.org/10.5194/bg-9-5023-2012>, 2012.
- Wesely, M.: Parameterization of surface resistances to gaseous dry deposition in regional-scale numerical models, *Atmos. Environ.*, 23, 1293-1304, 1989.
- Zaveri, R. A., and Peters, L. K.: A new lumped structure photochemical mechanism for large-scale applications, *J. Geophys. Res. Atmos.*, 104, 30387-30415, <https://doi.org/10.1029/1999JD900876>, 1999.
- Zaveri, R. A., Easter, R. C., Fast, J. D., and Peters, L. K.: Model for Simulating Aerosol Interactions and Chemistry (MOSAIC), *J. Geophys. Res. Atmos.*, 113, D13204, <https://doi.org/10.1029/2007JD008782>, 2008.

## **Response to Referee #2**

*Referee: The study analyses the possible factors responsible for elevated levels of ammonia over Indo-Gangetic Plain (IGP) regions during summertime. From previous studies, IGP has been identified as a hotspot for ammonia and analyzing the reasons for the high levels is an important study. Considering the implications of high-level ammonia, this study improves the understanding of the scientific community working on the ammonia over IGP. The authors use the air quality model (WRF-Chem), emission inventory and observation data (IASI satellite) to analyze the possible factors namely chemical conversion, emissions, and meteorology on ammonia loading. Specific comments with line numbers are provided below and I recommend these issues to be addressed before publication in ACP.*

**Response:** We would like to thank the referrer for your detailed and constructive comments. Please see our point-by-point reply below.

### ***Referee: 1. Title***

*The authors classify June to August as the summer period in the study. However, the Indian summertime is considered from April-June and monsoon begins from mid-June/late June. Summer to monsoon season has a drastic change in regional meteorology over IGP and possibility impact the levels of ammonia (gas phase) over IGP from high to low because of washout effects? I would suggest the authors not to mention the summer season in the title or in the study in general if possible or provide some explanation here.*

**Response:** Accepted. As you suggested, we have corrected “summer” to “pre-monsoon and monsoon seasons”. The title and the relevant part of main text have been revised.

### ***Referee: 2. Methods***

*Line 65: The authors mentioned the fair agreement of IASI ammonia observations with ground-based measurements citing few studies which are mainly satellite observations. It is not clear whether the fair agreement was for IGP or other regions. It would be helpful if the authors could provide some details about the ground measurements used here.*

**Response:** Accepted. The two articles cited here are validations of IASI NH<sub>3</sub> measurements using ground-based and airborne data sets around the world. In Dammers et al. (2016), IASI-NH<sub>3</sub> measurements were evaluated with ground-based Fourier transform infrared spectroscopy (FTIR) measurements from nine Network for the Detection of Atmospheric Composition Change (NDACC) stations around the world between 2008 and 2015, and results showed that IASI NH<sub>3</sub> were in fair agreement with ground-based measurements with a mean relative difference of  $-32.4 \pm 56.3\%$  and a correlation  $r$  of 0.8. In Van Damme et al. (2015), IASI-NH<sub>3</sub>



measurements were validated using existing independent ground-based and airborne data sets in Europe, China and Africa, results showed that IASI NH<sub>3</sub> were generally consistent with the available data sets. However, due to the lack of publicly available ammonia observation data sets in the IGP, previous studies have not evaluated IASI NH<sub>3</sub> in the IGP. For the same reason, the evaluation of IASI NH<sub>3</sub> in the IGP cannot be implemented in this study. As you suggested, we reworded in the revised manuscript.

**Revision:** (Page 3, Line 72-75) “It has been found that the IASI samples at the overpass time could represent the entire day, and IASI NH<sub>3</sub> observations are in fair agreement with the available ground-based and airborne data sets around the world (Dammers et al., 2016; Van Damme et al., 2015). However, due to the lack of publicly available ammonia observation data sets in the IGP, previous studies have not evaluated IASI NH<sub>3</sub> in the IGP.”

*Line 77: The authors should provide more details on the input of the ISORROPIA-II model or in general the model in order to interpret the gas-particle partitioning of the ammonia for example what are the gas species in input?*

**Response:** Accepted. As you suggested, we added more details about the inputs of ISORROPIA-II in Sect 2.3.

**Revision:** (Page 3, Line 92-94) “The chemical and meteorological data from WRF-Chem, including water-soluble ions (SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup>, Cl<sup>-</sup>, Na<sup>+</sup>) in PM<sub>2.5</sub>, gaseous precursors (NH<sub>3</sub>, HNO<sub>3</sub>, HCl), temperature (T) and relative humidity (RH) are used as the inputs of ISORROPIA-II.”

*Line 80: “As inputs of ISORROPIA-II, the outputs (water-soluble ions, gas species, T and RH) of WRF-Chem were first averaged over the IGP and then averaged for summer 2010”. Please provide more details on averaging here. As IGP is a vast region, averaging the inputs over the region may create a bias for some regions over IGP considering the heterogeneity of the sources over IGP. Again the time averaging may create some bias too as the time period included in the study has few days or a month of the summer season and two months of monsoon as per Indian meteorological department classification. It would be useful to check the variability of the inputs month wise for example June, July, and August separately.*

**Response:** Accepted. Previously, the offline calculation of ISORROPIA-II was based on the average of the entire IGP from June to August, which might cause bias due to the spatial and temporal heterogeneity. As you suggested, we divided the IGP from northwest to southeast into three areas (namely western IGP, central IGP, and eastern IGP), and divided the study period into two seasons (pre-monsoon and monsoon), and analyzed the sensitivity of NH<sub>3</sub> to temperature for the three regions under the two seasons in Sect 3.2.

**Revision:** (Page 4, Line 98-101) “Sensitivity tests was firstly conducted based on the average of the entire IGP from June to August. Then, the IGP was divided equally from northwest to southeast into three regions (namely western IGP, central IGP, and eastern IGP), and the study period was divided into the pre-monsoon season (June) and the monsoon season (July to August). Sensitivity tests were conducted for the three regions under the two seasons.”

(Page 6, Line 182-186) “As shown in Figure 3c, the sensitivity of  $\text{NH}_3$  to temperature varies in different seasons and regions. Temporally, the sensitivity of  $\text{NH}_3$  to temperature during the monsoon season is generally higher than that during the pre-monsoon season. Spatially, the sensitivity of  $\text{NH}_3$  to temperature is highest over the eastern IGP, followed by the central IGP and the western IGP. The difference in the sensitivity of the  $\text{NH}_3$  to temperature may be caused by the difference of the initial  $\varepsilon(\text{NH}_4^+)$  and temperature.”

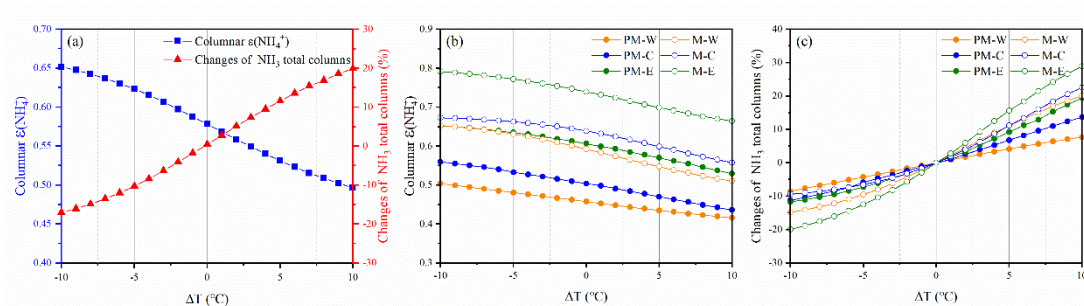


Figure 3. Columnar  $\varepsilon(\text{NH}_4^+)$  and changes of  $\text{NH}_3$  total columns with the changes of temperatures predicted by ISORROPIA-II. (a) Mean columnar  $\varepsilon(\text{NH}_4^+)$  and changes of  $\text{NH}_3$  total columns over the IGP from June to August 2010. (b) Columnar  $\varepsilon(\text{NH}_4^+)$  and (c) changes of  $\text{NH}_3$  total columns over the western IGP during Pre-monsoon (PM-W), the central IGP during Pre-monsoon (PM-C), the eastern IGP during pre-monsoon (PM-E), the western IGP during monsoon (M-W), the central IGP during monsoon (M-C), the eastern IGP during monsoon (M-E).

### Referee: 3. Results

Line 105: As per the sowing season, IGP has mainly two cropping cycles which includes summer and winter both. During both cycles, fertilizer applications can be intensive. Is there any study/data supporting the highest N-fertilizer application during the months mentioned in this study? Line108: Authors cited the Riddick et al. 2010 for the exponential increase of ammonia emission with temperature. Would it be possible to expect similar results, if the temperature data and ammonia satellite observations from the present study taken in to account?

**Response:** The monthly map of N-fertilizer application amounts from Nishina et al. (2017) shows that there two peaks in N-fertilizer application amounts over the IGP, one in May-August, the other in November-December, which is consistent with the two cropping cycles in the IGP. We reworded in the

revised manuscript.

However, it is difficult to study the relationship between NH<sub>3</sub> emission and temperature in this study. Because NH<sub>3</sub> satellite observations are affected by many factors, and are not linearly related with NH<sub>3</sub> emissions. Thus, it is difficult to isolate the influence of temperature using NH<sub>3</sub> observations.

**Revision:** (Page 4, Line 121-125) “IGP has two cropping cycles including summer and winter (2012). June to August is one of the two main sowing periods in the IGP with a large amount of N-fertilizer applied to the cropland as base fertilizer. The monthly map of N-fertilizer application amounts from Nishina et al. (2017) shows that there are two peaks in N-fertilizer application amounts over the IGP with one in May-August, the other in November-December, which is consistent with the two cropping cycles in the IGP.”

*Line 160: The analysis of the low gas-to-particle conversion of ammonia demonstrates the sensitivity to SO<sub>2</sub>/NO<sub>x</sub> emissions. This is an excellent analysis based on the model output data. Is it fair to expect similar results if the data for SO<sub>2</sub>/NO<sub>x</sub> not modeled but taken from observations either ground or satellite observations? Please comment on this.*

**Response:** Accepted. The spatial distributions of satellite-derived SO<sub>2</sub> and NO<sub>2</sub> columns, as shown in Figure S2, further demonstrate the lower SO<sub>2</sub> and NO<sub>2</sub> concentrations over the IGP compared to the NCP. The SO<sub>2</sub> and NO<sub>2</sub> columns over the IGP ( $0.5$  and  $2.3 \times 10^{15}$  molecules cm<sup>-2</sup>) are much lower than that over the NCP ( $10.4$  and  $8.3 \times 10^{15}$  molecules cm<sup>-2</sup>).

**Revision:** (Page 3, Line 78-81) “SO<sub>2</sub> columns from June to August 2010 were derived from the Level-3 Aura/OMI Global SO<sub>2</sub> Data Products (OMSO2e) (Krotkov et al., 2015). Tropospheric NO<sub>2</sub> columns from Ozone Monitoring Instrument (OMI) aboard NASA Aura satellite were used from June to August 2010 ([http://www.temis.nl/airpollution/no2col/no2regioomimonth\\_qa.php](http://www.temis.nl/airpollution/no2col/no2regioomimonth_qa.php))”

(Page 5, Line 136-138) “Besides, the satellite-derived SO<sub>2</sub> and NO<sub>2</sub> columns over the IGP ( $0.5$  and  $2.3 \times 10^{15}$  molecules cm<sup>-2</sup>) are also much lower than that over the NCP ( $10.4$  and  $8.3 \times 10^{15}$  molecules cm<sup>-2</sup>) (Fig. S2).”

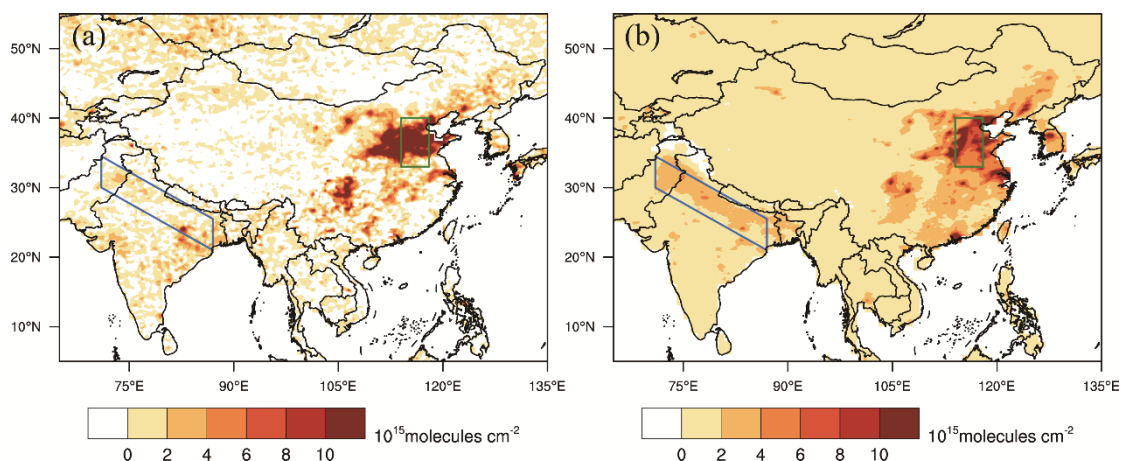


Figure S2. The spatial distribution of (a) SO<sub>2</sub> and (b) NO<sub>2</sub> columns over East Asia from June to August 2010.

*Line 200: The authors concluded an interesting observation about the other factors than wet conditions and high RH controlling the high ammonia loading over the IGP. More details on this would be useful to support this conclusion. Even a simple time series of IASI ammonia observations with RH for the period of study would provide an important context.*

**Response:** Accepted. As you suggested, we added more details about the effect of RH on gas-to-particle partitioning of NH<sub>3</sub> in Sect 3.3.

**Revision:** (Page 8, Line 225-231) “The simulated surface  $\epsilon(\text{NH}_4^+)$  over the western, central and eastern part of the IGP were 0.11, 0.13 and 0.24 during pre-monsoon and 0.26, 0.26 0.37 during monsoon. It is not difficult to find that the surface  $\epsilon(\text{NH}_4^+)$  during the monsoon season is significantly higher than that during the pre-monsoon season, and the surface  $\epsilon(\text{NH}_4^+)$  generally increases from northwest to southeast along the IGP. Besides, the columnar  $\epsilon(\text{NH}_4^+)$  shows similar spatiotemporal variations with the surface  $\epsilon(\text{NH}_4^+)$  (Figure 3b). The spatiotemporal variations of  $\epsilon(\text{NH}_4^+)$  are consistent with the spatiotemporal variations of RH (Figure S5a), indicating that RH is an important factor affecting the NH<sub>3</sub> partitioning.”

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# Why the Indo-Gangetic Plain is the region with the largest NH<sub>3</sub> column in the globe during **pre-monsoon and monsoon seasons**?

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**Abstract.** Satellite observations show a global maximum in ammonia (NH<sub>3</sub>) over the Indo-Gangetic Plain (IGP), with a peak **from June to August**. However, it has never been explained explicitly. In this study, we investigated the causes of high NH<sub>3</sub> loading over the IGP during **pre-monsoon and monsoon seasons** using WRF-Chem (Weather Research and Forecasting model coupled to chemistry). IGP has relatively high NH<sub>3</sub> emission fluxes (0.4 t km<sup>-2</sup> month<sup>-1</sup>) due to intensive agricultural activities and high air temperature **from June to August**. Additionally, low sulfur dioxide (SO<sub>2</sub>) and nitrogen oxides (NO<sub>x</sub>) emissions and high air temperature limit the gas-to-particle conversion of NH<sub>3</sub>, particularly for ammonium nitrate formation. Moreover, the barrier effects of the Himalayas in combination with the surface convergence weaken the horizontal diffusion of NH<sub>3</sub>. The high NH<sub>3</sub> loading over the IGP mainly results from the low gas-to-particle partitioning of NH<sub>3</sub> caused by low SO<sub>2</sub> and NO<sub>x</sub> emissions. It contrasts to those in the North China Plain, where high SO<sub>2</sub> and NO<sub>x</sub> emissions promote the conversion of gaseous NH<sub>3</sub> into particulate ammonium.

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## 1 Introduction

Ammonia (NH<sub>3</sub>) has multiple environmental implications. As the only alkaline gas in the atmosphere, it reacts with sulfuric acid (H<sub>2</sub>SO<sub>4</sub>) or nitric acid (HNO<sub>3</sub>) to produce ammonium (NH<sub>4</sub><sup>+</sup>) containing aerosols (Seinfeld and Pandis, 2006), which can affect Earth's radiative balance (Abbatt et al., 2006; Adams et al., 2001) and endanger public health (Pope et al., 2002; Stokstad, 2014). In addition, NH<sub>3</sub> is the main form of reactive nitrogen in the environment (Reis et al., 2009), the deposition of ammonia and ammonium can cause acidification of terrestrial ecosystems and eutrophication of water bodies (Paerl et al., 2014).

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Satellite observations (Van Damme et al., 2018; Warner et al., 2016) and ground-based measurements (Carmichael et al., 2003) have revealed that the Indo-Gangetic Plain (IGP) has the global maximum NH<sub>3</sub> loading, particularly **from June to August**. Previous studies have suggested that the high NH<sub>3</sub> loading over the IGP is caused by high NH<sub>3</sub> emissions from intensive agricultural activities (Clarisse et al., 2009; Van Damme et al., 2015b). Interestingly, satellite measurements show that the total columns of NH<sub>3</sub> over the IGP are much higher than those over the North China Plain (NCP), which has higher

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NH<sub>3</sub> emissions fluxes ([www.meicmodel.org/dataset-mix](http://www.meicmodel.org/dataset-mix)). Therefore, emissions alone might not be enough to explain the high NH<sub>3</sub> loading over the IGP.

Apart from dry deposition and wet removal by precipitation, another main sink for NH<sub>3</sub> is scavenging by acidic species to form particulate NH<sub>4</sub><sup>+</sup>. H<sub>2</sub>SO<sub>4</sub> and HNO<sub>3</sub> resulting from the oxidation of sulfur dioxide (SO<sub>2</sub>) and nitrogen oxides (NO<sub>x</sub>) are major acidic species in the atmosphere. Previous studies have confirmed that reduced SO<sub>2</sub> and NO<sub>x</sub> emissions are key factors driving the increase in NH<sub>3</sub> concentration (Liu et al., 2018; Yu et al., 2018; Warner et al., 2017). In addition, meteorological conditions (including wind speed, precipitation, relative humidity and air temperature) also influence NH<sub>3</sub> loading through various chemical and physical processes. These factors may be causing the high NH<sub>3</sub> loading over the IGP, but these assumptions have not been verified in a modeling study.

In this study, we use a regional air quality model to investigate the causes of high NH<sub>3</sub> loading over the IGP during **pre-monsoon and monsoon seasons**. This is the first study to analyze the causes of high NH<sub>3</sub> loading over the IGP considering all possible factors. The remainder of this paper is organized as follows. The air quality model and observational data are described in sect. 2. Section 3 analyzes the influences of several factors (including emissions, chemical conversion, and meteorological conditions) on NH<sub>3</sub> loading. Among them, SO<sub>2</sub> and NO<sub>x</sub> emissions over the IGP are compared to those over the NCP to clearly illustrate their impacts on NH<sub>3</sub> loading. Section 4 provides concluding remarks.

## 2 Methods

### 2.1 WRF-Chem model and emissions inventory

WRF-Chem (Fast et al., 2006; Grell et al., 2005) version 3.6.1 was applied to investigate the cause of the high NH<sub>3</sub> loading over the IGP during **pre-monsoon and monsoon seasons**. The simulation was performed on a domain with 30 km horizontal resolution covering the northern part of India and parts of Pakistan, Nepal, China, and Bangladesh with 120 × 90 grid cells. There were 23 vertical levels from the surface to the top pressure of 50 hPa. The simulations were conducted from 25 May to 31 August 2010 and the first 7 days (25-31 May) were treated as the spin-up period. **June was considered pre-monsoon season. July to August was considered monsoon season.** The initial meteorological and boundary conditions were obtained from the National Centers for Environmental Prediction Final Analysis with a 6 h temporal resolution. **CBM-Z (Carbon Bond Mechanism version Z) chemical mechanism (1) and MOSAIC (Model for Simulating Aerosol Interactions and Chemistry) aerosol module (2) were used for modeling photochemical reactions and aerosol processes (nucleation, growth and thermodynamic equilibrium), respectively. Wet deposition in the model includes both in-cloud and below-cloud scavenging. The blow-cloud scavenging of aerosols by impaction and trace gases by mass transfer was based on the methods of Easter et al. (3). Dry deposition was treated following the methods of Wesely (4) for trace gases, and Binkowski and Shankar (5) for aerosols.** More model configurations are described in Table S1.

Anthropogenic emissions were obtained from the MIX inventory (Li et al., 2017), an Asian anthropogenic emissions inventory that harmonizes several local inventories using a mosaic approach. MIX uses Regional Emissions Inventory in Asia (REAS2, version 2) (Kurokawa et al., 2013) for NH<sub>3</sub> emissions in India.

## 2.2 Observational dataset

Atmospheric total columns of NH<sub>3</sub> were derived from measurements of an Infrared Atmospheric Sounding Interferometer (IASI) on board MetOp-A (<https://iasi.aeris-data.fr/NH3/>). Metop-A was launched in 2006 in a Sun-synchronous orbit with a mean local solar overpass time of 9:30 a.m. and 9:30 p.m. Only the daytime measurements have been used here, because the nighttime measurements had larger relative errors caused by the general lower thermal contrast for the nighttime overpass (Van Damme et al., 2014). It has been found that the IASI samples at the overpass time could represent the entire day, and IASI NH<sub>3</sub> observations are in fair agreement with the available ground-based and airborne data sets around the world (Dammers et al., 2016; Van Damme et al., 2015a). However, due to the lack of publicly available ammonia observation data sets in the IGP, previous studies have not evaluated IASI NH<sub>3</sub> in the IGP. This work used the ANNI-NH3-v2.2R-I retrieval product, which relied on ERA-Interim reanalysis for its meteorological inputs (Van Damme et al., 2017). The mean NH<sub>3</sub> column concentrations over East Asia on a 0.25° × 0.25° grid from June to August 2010 have been determined based on the relative error weighting mean method (Van Damme et al., 2014). SO<sub>2</sub> columns from June to August 2010 were derived from the Level-3 Aura/OMI Global SO<sub>2</sub> Data Products (OMSO2e) (Krotkov et al., 2015). Tropospheric NO<sub>2</sub> columns from Ozone Monitoring Instrument (OMI) aboard NASA Aura satellite were used from June to August 2010 ([http://www.temis.nl/airpollution/no2col/no2regioomimonth\\_qa.php](http://www.temis.nl/airpollution/no2col/no2regioomimonth_qa.php)).

Meteorological data at 38 sites over northern India obtained from the National Climate Data Center (<https://gis.ncdc.noaa.gov/maps/ncei/cdo/hourly>) were used to evaluate the accuracy of meteorological simulations. The evaluated variables included hourly wind speed at 10 m (WS10), wind direction at 10 m (WD10), relative humidity at 2 m (RH2) and temperature at 2 m (T2). The statistical parameters included the mean bias (MB), normalized mean bias (NMB), root mean square error (RMSE) and correlation coefficient (R). In addition, air temperature at 21 sites over the NCP obtained from the National Climate Data Center were also used in this work.

## 2.3 ISORROPIA-II thermodynamic model

The thermodynamic equilibrium model, ISORROPIA-II (Fountoukis and Nenes, 2007), treating the thermodynamics of NH<sub>4</sub><sup>+</sup>-SO<sub>4</sub><sup>2-</sup>-NO<sub>3</sub><sup>-</sup>-K<sup>+</sup>-Ca<sup>2+</sup>-Mg<sup>2+</sup>-Na<sup>+</sup>-Cl<sup>-</sup>-H<sub>2</sub>O aerosol system, was used to investigate the influence of air temperature on the NH<sub>3</sub> total columns. In this study, ISORROPIA-II was run in the “forward mode” and assuming particles are “metastable” with no solid precipitates. The chemical and meteorological data from WRF-Chem, including water-soluble ions (SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup>, Cl<sup>-</sup>, Na<sup>+</sup>) in PM<sub>2.5</sub>, gaseous precursors (NH<sub>3</sub>, HNO<sub>3</sub>, HCl), temperature (T) and relative humidity (RH) are used as the inputs of ISORROPIA-II. Using ISORROPIA-II, we simulated 20 scenarios. In these cases, air temperature of each layer increased or decreased by 10 °C synchronously, with the interval of 1 °C. Meanwhile, the other input parameters



remained the same. Then, we calculated the columnar  $\epsilon(\text{NH}_4^+)$  (partitioning ratios of  $\text{NH}_4^+$  to total ammonia (TA,  $\text{TA} = \text{NH}_3 + \text{NH}_4^+$ )) in each case. The columnar  $\epsilon(\text{NH}_4^+)$  is the sum of the  $\epsilon(\text{NH}_4^+)$  of each vertical level, but each weighted by the thickness of the layer and mass concentration of TA. Sensitivity tests was firstly conducted based on the average of the entire IGP from June to August. Then, the IGP was divided equally from northwest to southeast into three regions (namely western IGP, central IGP, and eastern IGP), and the study period was divided into the pre-monsoon season (June) and the monsoon season (July to August). Sensitivity tests were conducted for the three regions under the two seasons.

### 3 Results

#### 3.1 High $\text{NH}_3$ emissions

IGP is a vast stretch of fertile alluvial plain spanning the banks of the Indus and Ganges Rivers and their tributaries. The main part of the IGP is located in India. The estimated  $\text{NH}_3$  emissions in India in 2010 was 9.9 Tg, which is comparable to that in China (9.8 Tg) and accounts for about 34 % of total  $\text{NH}_3$  emissions in Asia (Li et al., 2017). Agriculture is the largest  $\text{NH}_3$  emitter in India, accounting for about 76 % of the total  $\text{NH}_3$  emissions (Li et al., 2017). Agricultural  $\text{NH}_3$  emissions mainly originate from animal husbandry and fertilizer application (Bouwman et al., 1997; Streets et al., 2003). India is the second largest N-fertilizer consumer (after China) and consumes 16.5 Tg N-fertilizers (16 % of the world's total) (FAOSTAT, 2010). In addition, there are an estimated 302 million cattle and buffalo in India (19 % of world's total), which is more than any other country (FAOSTAT, 2010). It is estimated that cattle and buffalo account for about 80 % of  $\text{NH}_3$  emissions among livestock in India (Aneja et al., 2012).

$\text{NH}_3$  emissions over the IGP was 4.3 Tg in 2010 (estimated using the MIX database), which was mainly attributed to intensive agricultural practices. The IGP is known as the food bowl of India spreading across the states of Punjab, Haryana, Uttar Pradesh, Bihar, and West Bengal (blue quadrangle in Fig. 1). The total number of cattle and buffalo in the five states was estimated to be 103 million (34 % of the national total) in 2012 (2012a). The total amount of N-fertilizer applied in the five states was estimated to be 6.9 Tg (42 % of the national total) in 2010 (2012b).  $\text{NH}_3$  emissions over the IGP from June to August are very high with a regional mean  $\text{NH}_3$  emissions flux of  $0.4 \text{ t km}^{-2} \text{ month}^{-1}$  (estimated using MIX database for 2010). This is consistent with satellite observations, which also show a peak of  $\text{NH}_3$  columns over the IGP from June to August (Van Damme et al., 2015b). The peak of  $\text{NH}_3$  emissions over the IGP might be the joint result of intensive N-fertilizer applications and high temperature. IGP has two cropping cycles including summer and winter (2012b). June to August is one of the two main sowing periods in the IGP with a large amount of N-fertilizer applied to the cropland as base fertilizer. The monthly map of N-fertilizer application amounts from Nishina et al. (2017) shows that there are two peaks in N-fertilizer application amounts over the IGP with one in May-August, the other in November-December, which is consistent with the two cropping cycles in the IGP. In addition, the air temperature is very high over the IGP with an observed regional mean value of  $30.9 \text{ }^\circ\text{C}$  from June to August 2010. Ammonia emissions increase exponentially with

temperature (Riddick et al., 2016). The high application rate of N-fertilizer and high air temperature could cause high NH<sub>3</sub> emissions, resulting in the high NH<sub>3</sub> columns.

130 The spatial distribution of mean NH<sub>3</sub> total columns over East Asia from June to August 2010 is shown in Fig. 1. The NH<sub>3</sub> columns over the IGP ( $7.6 \times 10^{16}$  molecules cm<sup>-2</sup>) were about twice as large as what was observed over the NCP ( $4.1 \times 10^{16}$  molecules cm<sup>-2</sup>). The NCP is also a large agricultural region (Huang et al., 2012). The regional mean NH<sub>3</sub> emissions flux over the NCP was 0.7 t km<sup>-2</sup> month<sup>-1</sup> from June to August 2010 (estimated using the MIX database), which was about 1.8 times that of the IGP. The IGP has much higher NH<sub>3</sub> total columns (Fig. 1) compared to the NCP, but lower NH<sub>3</sub> emissions fluxes (Fig. S1a). Therefore, other factors might lead to the high NH<sub>3</sub> loading over the IGP besides high NH<sub>3</sub> emissions.

### 135 3.2 Low gas-to-particle conversion of NH<sub>3</sub>

The emissions fluxes of SO<sub>2</sub> and NO<sub>x</sub> (both are 0.3 t km<sup>-2</sup> month<sup>-1</sup>) over the IGP are only about one-fourth of that over the NCP (1.1 and 1.3 t km<sup>-2</sup> month<sup>-1</sup>) (Table 1 and Fig. S1). Besides, the satellite-derived SO<sub>2</sub> and NO<sub>2</sub> columns over the IGP (0.5 and  $2.3 \times 10^{15}$  molecules cm<sup>-2</sup>) are also much lower than that over the NCP ( $10.4$  and  $8.3 \times 10^{15}$  molecules cm<sup>-2</sup>) (Fig. S2). The relatively low SO<sub>2</sub> and NO<sub>x</sub> emissions could be an important factor causing the high NH<sub>3</sub> columns over the IGP. In this study, we used the molar ratio ( $R_{emis}$ ) of NH<sub>3</sub> emissions fluxes ( $E_A$ ) to the sum of twice the SO<sub>2</sub> emissions fluxes ( $E_S$ ) and NO<sub>x</sub> emissions fluxes ( $E_N$ ) to roughly represent the richness of NH<sub>3</sub> in the atmosphere, given by Eq. (1):

$$R_{emis} = \frac{E_A}{2 \times E_S + E_N} \quad (1)$$

The calculated  $R_{emis}$  in the IGP was 1.35, which was about 2.6 times as large as that in the NCP (0.51). We performed simulations for a base case and a ‘increased SO<sub>2</sub>/NO<sub>x</sub> emissions’ case to investigate the impact of SO<sub>2</sub> and NO<sub>x</sub> emissions on NH<sub>3</sub> loading. In the increased SO<sub>2</sub>/NO<sub>x</sub> emissions case, the emissions of SO<sub>2</sub> and NO<sub>x</sub> increased 2.6 times to make  $R_{emis}$  of the IGP equal to that of the NCP.

The simulated NH<sub>3</sub> columns in the base case are shown in Fig. 2a. It is noted that the IASI NH<sub>3</sub> columns cannot be quantitatively compared to modeled NH<sub>3</sub> columns as the IASI NH<sub>3</sub> products do not provide information on the vertical sensitivity (averaging kernels) to properly weight the model values. Nonetheless, the simulated regional mean NH<sub>3</sub> total column over the IGP from the base case of  $8.8 \times 10^{16}$  molecules cm<sup>-2</sup> is close to the satellite-derived value ( $7.6 \times 10^{16}$  molecules cm<sup>-2</sup>), indicating that the model could generally capture the magnitude of NH<sub>3</sub> columns. Additionally, a broadly similar pattern was found in the NH<sub>3</sub> columns in the base run as in the satellite observations, both of which showed that NH<sub>3</sub> columns decrease along the IGP from northwest to southeast with the highest values in northwestern IGP (Figs. 1 and 2a). The statistical performance of the meteorological predictions at 38 sites over Northern India are presented in Table S2. The predicted T2 matched well with the observations with a correlation coefficient of 0.8 and an NMB of 4.2 %. The predicted RH2 was slightly underestimated with an NMB of -13.4 % and a correlation coefficient of 0.8. The predicted WS10 agreed reasonably well with the observations with an NMB of -5.3 %. In addition, the simulated WD10 matched well with the

observations, and both the predicted and observed dominant wind direction was SSE. The good agreement between the simulation and the observations confirms the reliability of the meteorological prediction over the simulation domain.

160 The spatial distribution of the  $\text{NH}_3$  total column in the increased emissions case is shown in Fig. 2b. The  $\text{NH}_3$  total columns significantly decreased over the entire IGP, with a regional mean value of  $2.5 \times 10^{16}$  molecules  $\text{cm}^{-2}$  (a 72.2 % decrease compared to the base case). The surface  $\epsilon(\text{NH}_4^+)$  in the base case and the increased  $\text{SO}_2/\text{NO}_x$  emissions case are shown in Fig. 2 (panels c and d, respectively). The surface  $\epsilon(\text{NH}_4^+)$  in the base case was fairly low with a regional mean value of 0.3 over the IGP. **The simulated  $\epsilon(\text{NH}_4^+)$  in the 2010 monsoon in Delhi was 0.38, which is close to the observed  $\epsilon(\text{NH}_4^+)$  (0.39 in the**  
165 **2011 monsoon season in Delhi) (Singh and Kulshrestha, 2012).** In the increased  $\text{SO}_2/\text{NO}_x$  emissions case, the regional mean surface  $\epsilon(\text{NH}_4^+)$  increased to 0.6 over the IGP. Significant increases in surface  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  concentrations were also found (Fig. S3). The regional mean surface  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  concentrations increased from 9.7 to 24.9 and from 7.2 to 20.0  $\mu\text{g m}^{-3}$ , respectively. Additionally, the regional mean columnar  $\epsilon(\text{NH}_4^+)$  over the IGP is 0.56 in the base case and increases to 0.87 in the increased  $\text{SO}_2/\text{NO}_x$  emissions case. This suggests that the increased  $\text{SO}_2$  and  $\text{NO}_x$  emissions enhanced the  
170 formation of acidic species and promoted the conversion of  $\text{NH}_3$  into  $\text{NH}_4^+$ . The effectively reduced  $\text{NH}_3$  total columns in the increased  $\text{SO}_2/\text{NO}_x$  emissions case indicate that low  $\text{SO}_2$  and  $\text{NO}_x$  emissions could be the major cause of the high  $\text{NH}_3$  loading over the IGP.

Besides the amount of acidic species, air temperature is also an important factor affecting the thermodynamic equilibrium of  $\text{NH}_3$  between the gas phase and the particle phase. Higher air temperature limits the gas-to-particle conversion of  $\text{NH}_3$  and  
175 enhances volatilization of  $\text{NH}_4\text{NO}_3$  (Seinfeld and Pandis, 2006). The observed regional mean air temperature over the IGP **from June to August 2010 was 30.9 °C, about 4.9 °C higher than the NCP (26.0 °C).** As shown in Fig. 3a, the columnar  $\epsilon(\text{NH}_4^+)$  increases as temperature decreases. A 10°C decrease in temperature results in a 0.07 increase in  $\epsilon(\text{NH}_4^+)$  and a consequent 17 % decrease in  $\text{NH}_3$  total columns. Additionally, a 10°C increase in temperature results in a 0.08 decrease in  $\epsilon(\text{NH}_4^+)$  and a consequent 20 % increase in  $\text{NH}_3$  total columns. If the temperature over the IGP drops to the temperature  
180 typical of the NCP (a 4.9 °C decrease), the  $\text{NH}_3$  total columns over the IGP will only decrease by 10 %. In contrast, if the  $\text{SO}_2/\text{NO}_x$  emissions over the IGP increase to make the  $R_{\text{emis}}$  of the IGP equal to that of the NCP, the  $\text{NH}_3$  column over the IGP will decrease by 72.2 %. Therefore, the low  $\text{SO}_2/\text{NO}_x$  emissions have a greater effect on causing high  $\text{NH}_3$  columns over the IGP than the high air temperature. **As shown in Figure 3c, the sensitivity of  $\text{NH}_3$  to temperature varies in different seasons and regions. Temporally, the sensitivity of  $\text{NH}_3$  to temperature during the monsoon season is generally higher than that during the pre-monsoon season. Spatially, the sensitivity of  $\text{NH}_3$  to temperature is highest over the eastern IGP, followed by the central IGP and the western IGP. The difference in the sensitivity of the  $\text{NH}_3$  to temperature may be caused by the difference of the initial  $\epsilon(\text{NH}_4^+)$  and temperature.**

### 3.3 Weak horizontal diffusion of $\text{NH}_3$

The IGP is located in the northern part of the Indian subcontinent and is surrounded by unique topography with the  
190 Himalayan range to the north and the Sulaiman range to the west. Weather on the Indian subcontinent is controlled by the

low-level Indian monsoon regime from June through September (Lawrence and Lelieveld, 2010). Fig. 4a shows the spatial distributions of surface wind flow and wind speed from June to August 2010. The dominant wind direction is southwest over the Indian peninsula and southeast over the IGP. Air mainly flows from the west coast of India and the south coast of Bengal. Surface wind speed is high on the west coast of India ( $>5 \text{ m s}^{-1}$ ) and on the south coast of Bengal ( $>4 \text{ m s}^{-1}$ ) but decreases from the coast inland. Mountains serve as barriers to the airflow on the surface of the Earth (Barry, 2008). Chow et al. (2013) reported that when stably stratified airflow encounters an extra-tropical mountain barrier, it is forced to rise and cool adiabatically. Consequently, higher pressure along the slope could be created, which could decelerate and block the flow. After a while, geostrophic adjustment occurs. As a result, the airflow turns left (right) in the northern (southern) hemisphere, and a barrier jet blowing parallel to the barrier is formed. As shown in Fig. 4a, the southerly airflow from the Bay of Bengal turns left when approaching the Himalayas, and then an easterly barrier jet parallel to the Himalayas is formed. The southwesterly airflow from the west coast of India also turns left when approaching the Himalayas. Similarly, wind flow at 850 hPa (Fig. 4b) also shows left-turning airflow near the Himalayas. The left-turning airflow indicates that the barrier effect of the Himalayas limits the northward movement of polluted air. Both satellite observations and the model simulation show that the high  $\text{NH}_3$  columns over the IGP are effectively cut off by mountains to the north (Fig. 2a and Fig. S4).

As shown in Fig. 4b, an area of low geopotential height extends from Pakistan to east India following the IGP. This elongated region of low pressure is known as the monsoon trough (Bohlinger et al., 2017). It causes wind to converge over this region. The convergence of horizontal wind can be observed from wind flow at both the surface and at 850 hPa. The prevailing wind directions south and east of the IGP are southwest and southeast, respectively. As a result of convergence of horizontal winds, an area of low wind speed forms and covers most of the IGP. The regional mean surface wind speed over the IGP is  $<3 \text{ m s}^{-1}$ . The weak wind speed in association with the convergence weakens the horizontal advection of  $\text{NH}_3$  and results in the accumulation of  $\text{NH}_3$  over the IGP.

The ventilation rate ( $V_r$ ) of the four edges of the IGP was used to illustrate the accumulation of an air mass over the IGP (Fig. 4d). The  $V_r$  of one edge is defined as the product of sectional area to the transport wind, given by Eq. (2):

$$V_r = AU_T \quad (2)$$

The sectional area A can be expressed as  $A = ZL$ , where Z is the mean boundary layer height along the edge and L is the

length of the edge. The transport wind  $U_T$  is given by 
$$U_T = \frac{1}{m} \sum_{j=1}^m \left( \frac{1}{n} \sum_{i=1}^n U_{ij} \right)$$
. m and n are the number of locations along the edge and vertical levels within the ABL where the winds are measured or predicted.  $U_{ij}$  is the wind speed perpendicular to the cross-section at each height and location along the edge. The ventilation rates of the four edges of the IGP were calculated using the WRF-Chem simulation results. The total  $V_r$  of the inflow from the southern and eastern edges ( $3.1 \times 10^9 \text{ m}^3 \text{ s}^{-1}$ ) was 64 % higher than the total  $V_r$  of the outflow from the western and northern edges ( $1.9 \times 10^9 \text{ m}^3 \text{ s}^{-1}$ ). The strong

inflow and weak outflow indicate accumulation of the air mass over the IGP. Therefore, outward transport of  $\text{NH}_3$  from the IGP through horizontal advection could be weak.

225 Interestingly, both relative humidity and precipitation are high over the IGP (Figs. S5), with regional mean values of 63 % and 660 mm from June to August 2010. The high relative humidity and precipitation suggest strong gas-to-particle conversion and wet scavenging of  $\text{NH}_3$  (Seinfeld & Pandis, 2006). The observed high  $\text{NH}_3$  loading under such a wet condition further indicates the effectiveness of other factors leading to high  $\text{NH}_3$  loading. The simulated surface  $\epsilon(\text{NH}_4^+)$  over the western, central and eastern part of the IGP were 0.11, 0.13 and 0.24 during pre-monsoon and 0.26, 0.26 0.37 during monsoon. It is not difficult to find that the surface  $\epsilon(\text{NH}_4^+)$  during the monsoon season is significantly higher than that during the pre-monsoon season, and the surface  $\epsilon(\text{NH}_4^+)$  generally increases from northwest to southeast along the IGP. Besides, the columnar  $\epsilon(\text{NH}_4^+)$  shows similar spatiotemporal variations with the surface  $\epsilon(\text{NH}_4^+)$  (Figure 3b). The spatiotemporal variations of  $\epsilon(\text{NH}_4^+)$  are consistent with the spatiotemporal variations of RH (Figure S5a), indicating that RH is an important factor affecting the  $\text{NH}_3$  partitioning. The meteorological conditions in the northwest IGP are characterized by higher air temperature, lower humidity, and lower rainfall compared to the southeast IGP (Figs. 4c and S5), all of which are conducive to the increase of  $\text{NH}_3$ . Consistently,  $\text{NH}_3$  total columns decrease from northwest to southeast along the IGP as revealed by both the satellite measurements and model simulations (Figs. 1 and 2a). However, emission fluxes of  $\text{NH}_3$  over the northwest IGP are also obviously higher than the southeast IGP (Fig. S1). To exclude the impact of emissions on the spatial distributions of  $\text{NH}_3$ , simulations for a “homogeneous emissions” case was performed by using WRF-Chem, where emissions of all primary pollutants over the IGP were set to their regional mean values. As shown in Fig. 5,  $\text{NH}_3$  total columns in the homogeneous emissions case still appear to decrease from northwest to southeast along the IGP. It is indicated that the meteorological factors (atmospheric diffusion, temperature, relative humidity, and precipitation) are important causes of the higher  $\text{NH}_3$  loadings over the northwest IGP than the southeast IGP.

#### 4 Conclusions

Satellite observations have revealed that the IGP has the global maximum  $\text{NH}_3$  loading with a peak from June to August. Our study reveals that the high  $\text{NH}_3$  loading over the IGP appears to be the joint result of high  $\text{NH}_3$  emissions, weak chemical loss, and weak horizontal diffusion. Intensive agricultural activities in combination with high temperature resulted in relatively high  $\text{NH}_3$  emissions over the IGP, with a regional mean  $\text{NH}_3$  emissions flux of  $0.4 \text{ t km}^{-2} \text{ month}^{-1}$ . The low  $\text{SO}_2$  and  $\text{NO}_x$  emissions and high temperature limited the conversion of  $\text{NH}_3$  to  $\text{NH}_4^+$ . The low chemical loss of gaseous  $\text{NH}_3$  is a key reason for the high  $\text{NH}_3$  loading over the IGP. In addition, orographic and meteorological conditions also play important roles in  $\text{NH}_3$  accumulation over the IGP. The barrier effects of the Himalayas limit the northward movement of monsoon air. The low wind speed ( $<3 \text{ m s}^{-1}$ ) in association with the surface convergence over the IGP weakens horizontal diffusion, which is conducive to the accumulation of  $\text{NH}_3$  over the IGP.

## Data availability

The IASI data used in this study was provided by the AERIS data infrastructure (<https://iasi.aeris-data.fr/NH3/>). The meteorological data used in this study was obtained from the National Climate Data Center integrated surface database (255 <https://gis.ncdc.noaa.gov/maps/ncei/cdo/hourly>). The anthropogenic emissions are available from MIX inventory ([www.meicmodel.org/dataset-mix](http://www.meicmodel.org/dataset-mix)). The SO<sub>2</sub> columns were provided by the NASA Goddard Earth Sciences Data and Information Services Center ([https://disc.gsfc.nasa.gov/datasets/OMSO2e\\_003/summary](https://disc.gsfc.nasa.gov/datasets/OMSO2e_003/summary)). The NO<sub>2</sub> columns are available from the Tropospheric Emission Monitoring Internet Service (260 [http://www.temis.nl/airpollution/no2col/no2regioomimonth\\_qa.php](http://www.temis.nl/airpollution/no2col/no2regioomimonth_qa.php)).

## Author contribution

Y.S initiated the investigation. T.W performed the modelling analyses. T.W, Y.S, Z.X and T.Z wrote and edited the manuscript. M.L, T.X, W.L, L.Y, X.C, H.Z and L.K contributed to discussions of the results and the manuscript.

## Competing interests

265 The authors declare no competing interests.

## Acknowledgements

This work was supported by the National Natural Science Foundation of China (NSFC) (91644212, 41675142 and 91837209) and the National Key R&D Program of China (2016YFC0201505).

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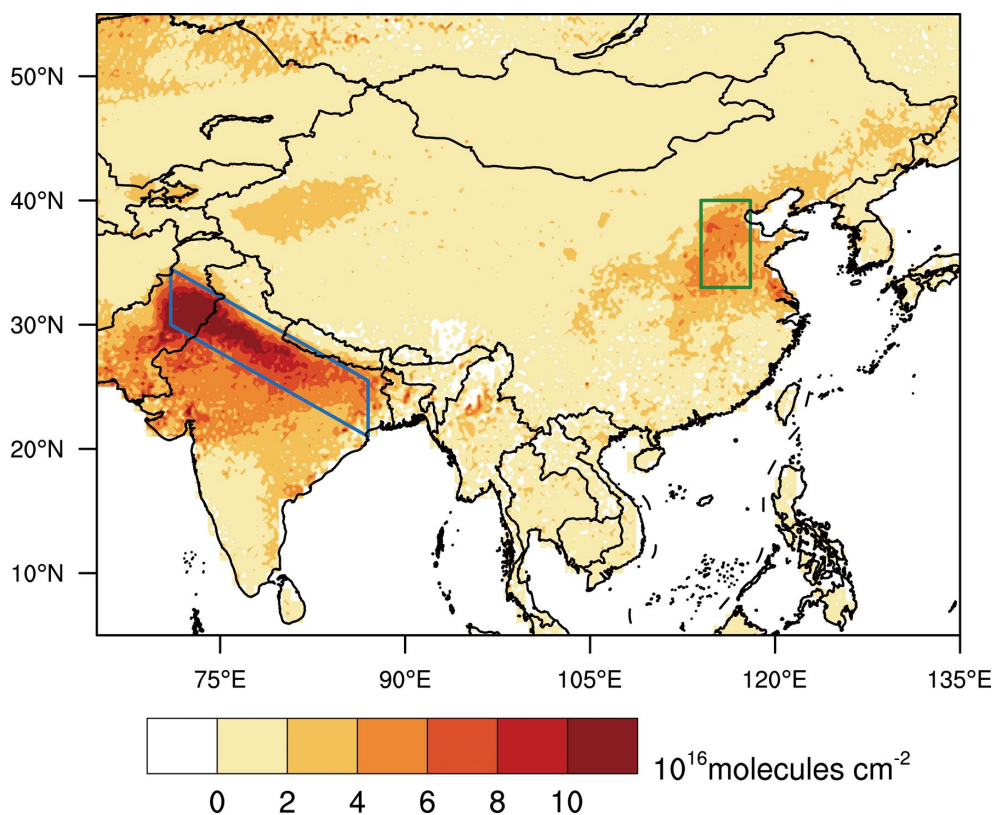


**Table 1.** Regional Mean NH<sub>3</sub> Total Columns and Emissions Fluxes of NH<sub>3</sub>, SO<sub>2</sub>, and NO<sub>x</sub> of the IGP and the NCP from June to August 2010.

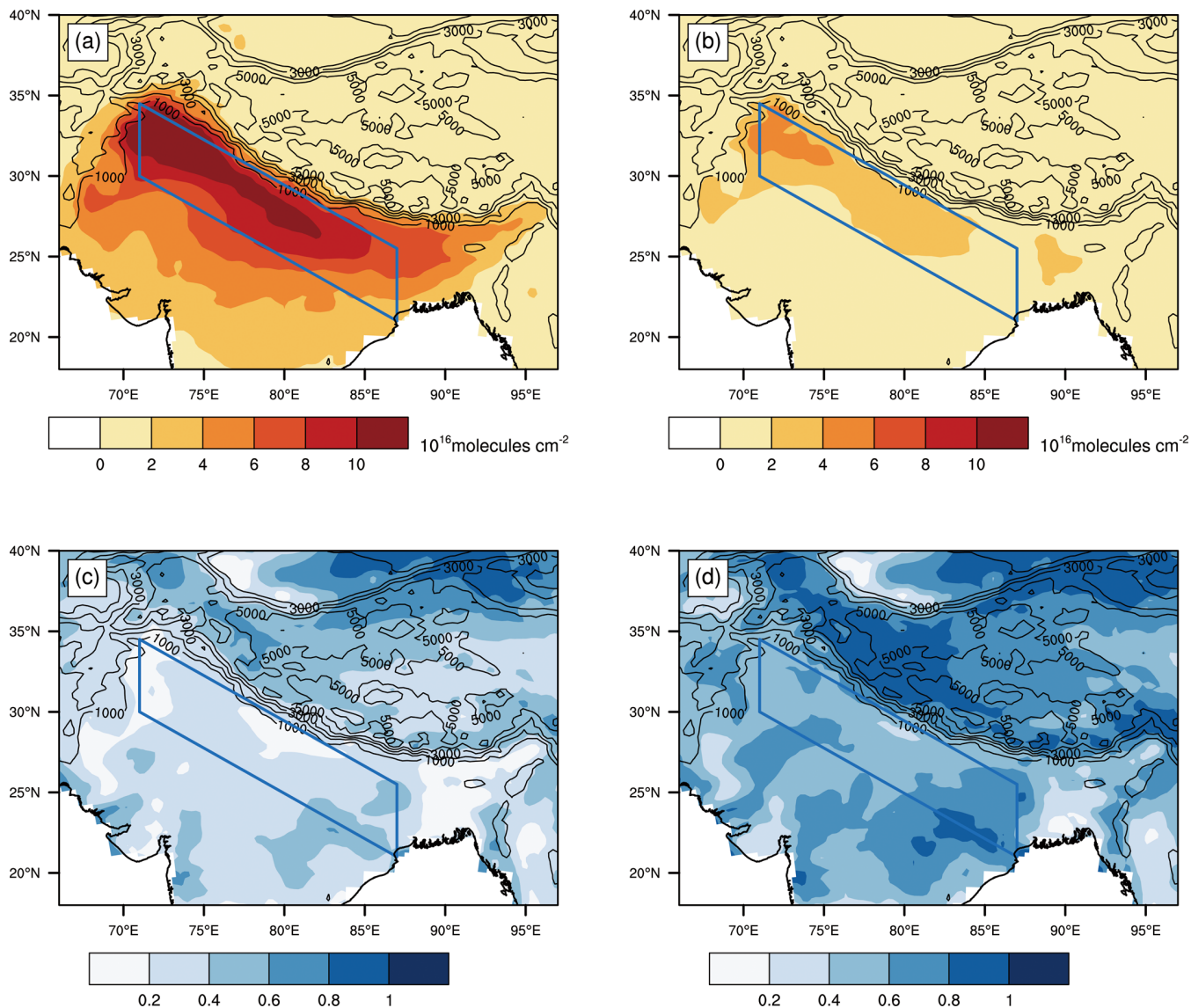
	NH <sub>3</sub> total columns <sup>a</sup> (molecules cm <sup>-2</sup> )	Emissions fluxes <sup>b</sup> (t km <sup>-2</sup> month <sup>-1</sup> )		
		NH <sub>3</sub>	SO <sub>2</sub>	NO <sub>x</sub>
IGP	7.6×10 <sup>16</sup>	0.4	0.3	0.3
NCP	4.1×10 <sup>16</sup>	0.7	1.1	1.3

370 <sup>a</sup>NH<sub>3</sub> total columns were derived from IASI measurements

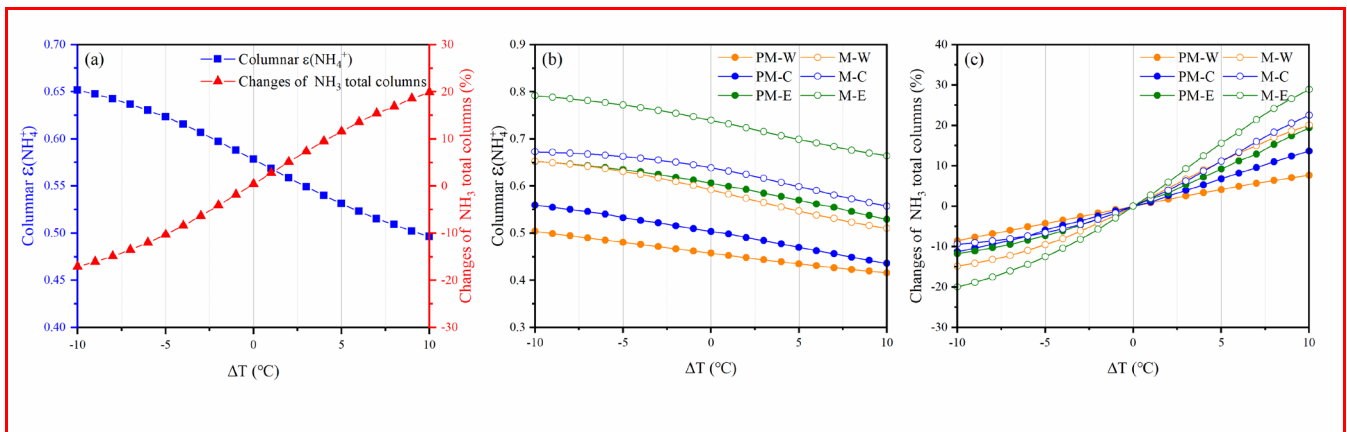
<sup>b</sup>Emissions fluxes were estimated using the MIX database



**Figure 1.** The spatial distribution of NH<sub>3</sub> total columns over East Asia from June to August 2010 retrieved from IASI measurements. The blue quadrangle represents the Indo-Gangetic Plain (IGP), and the green quadrangle represents the Northern China Plain (NCP).  
375

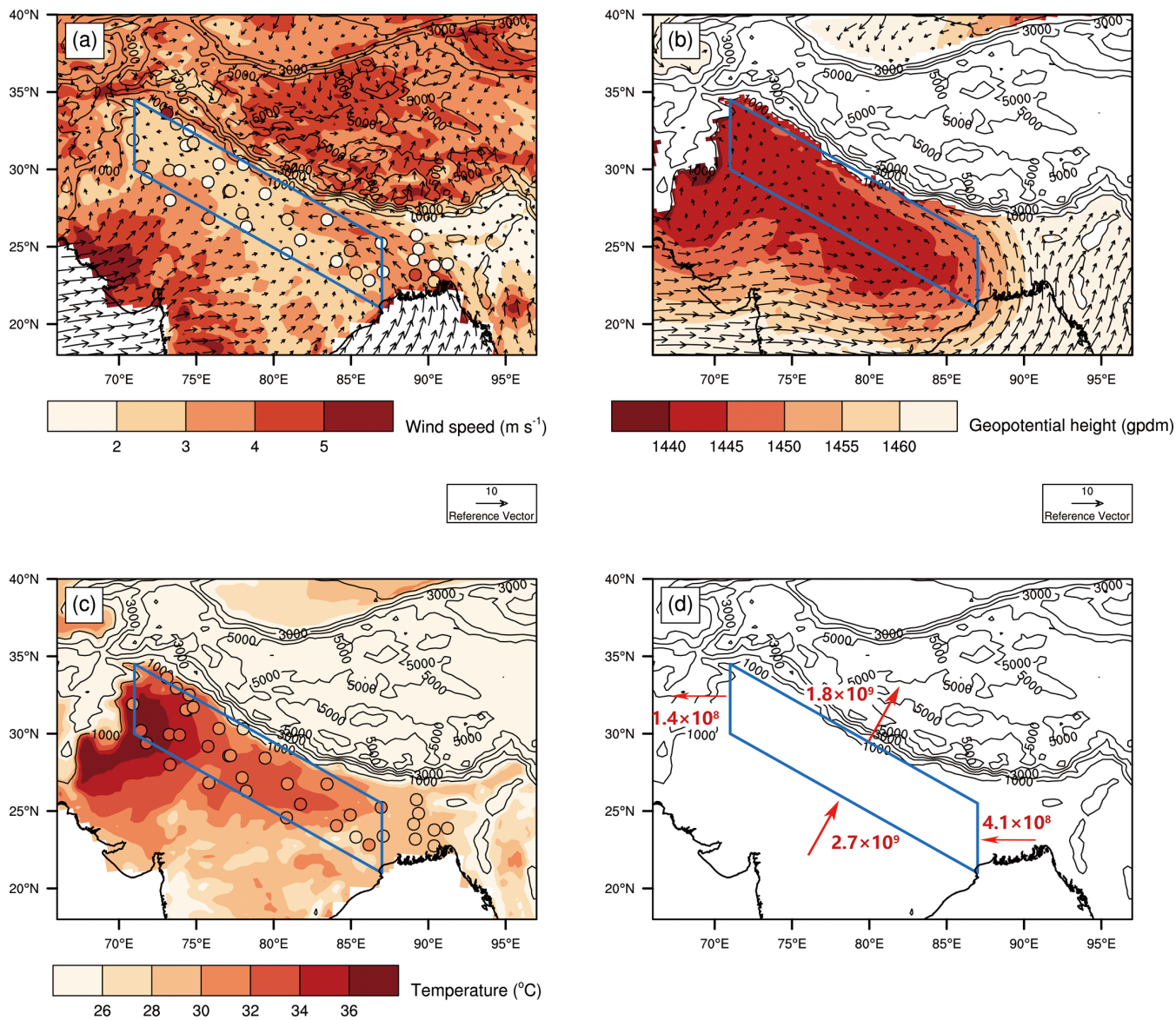


**Figure 2.** Spatial distributions of WRF-Chem predicted total columns of  $\text{NH}_3$  and surface  $\epsilon(\text{NH}_4^+)$  from June to August 2010. (a) and (b) are total columns of  $\text{NH}_3$  in the base case and the increased  $\text{SO}_2/\text{NO}_x$  emissions case, respectively. (c) and (d) are surface  $\epsilon(\text{NH}_4^+)$  in the base case and the increased  $\text{SO}_2/\text{NO}_x$  emissions case, respectively.

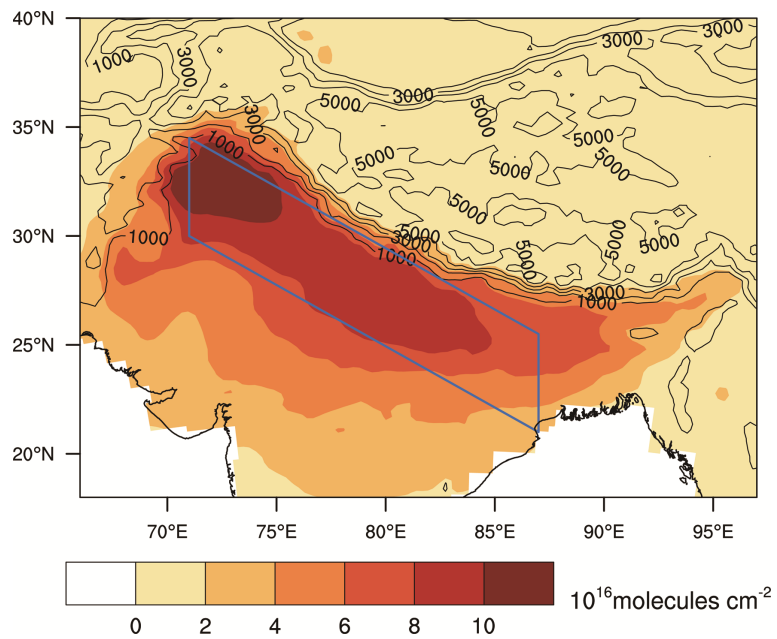


385 **Figure 3.** Columnar  $\epsilon(\text{NH}_4^+)$  and changes of  $\text{NH}_3$  total columns with the changes of temperatures predicted by ISORROPIA-II. (a) Mean columnar  $\epsilon(\text{NH}_4^+)$  and changes of  $\text{NH}_3$  total columns over the IGP from June to August 2010. (b) Columnar  $\epsilon(\text{NH}_4^+)$  and (c) changes of  $\text{NH}_3$  total columns over the western IGP during Pre-monsoon (PM-W), the central IGP during Pre-monsoon (PM-C), the eastern IGP during pre-monsoon (PM-E), the western IGP during monsoon (M-W), the central IGP during monsoon (M-C), the eastern IGP during monsoon (M-E).

390



**Figure 4.** Spatial distributions of WRF-Chem predicted meteorological variables from June to August 2010. (a) Wind flow and wind speed at 10 m. (b) Wind flow and geopotential height at 850 hPa. (c) Air temperature at 2 m. (d) Ventilation rate ( $\text{m}^3 \text{s}^{-1}$ ) of the four edges of the IGP. Circles in (a) and (c) show the observed wind speed at 10 m and air temperature at 2 m, respectively.



**Figure 5.** Spatial distributions of WRF-Chem predicted total columns of NH<sub>3</sub> from June to August 2010 in the homogeneous emissions case.

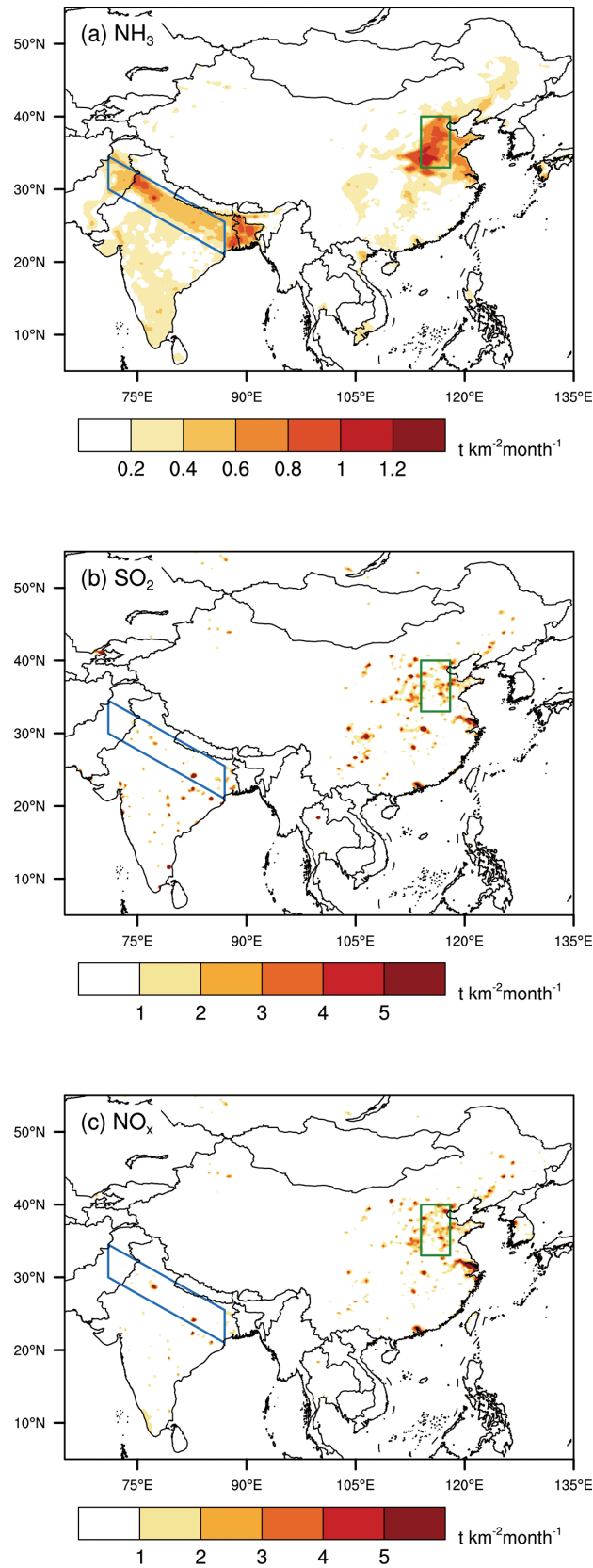
## Contents of this file

Figures S1 to S5

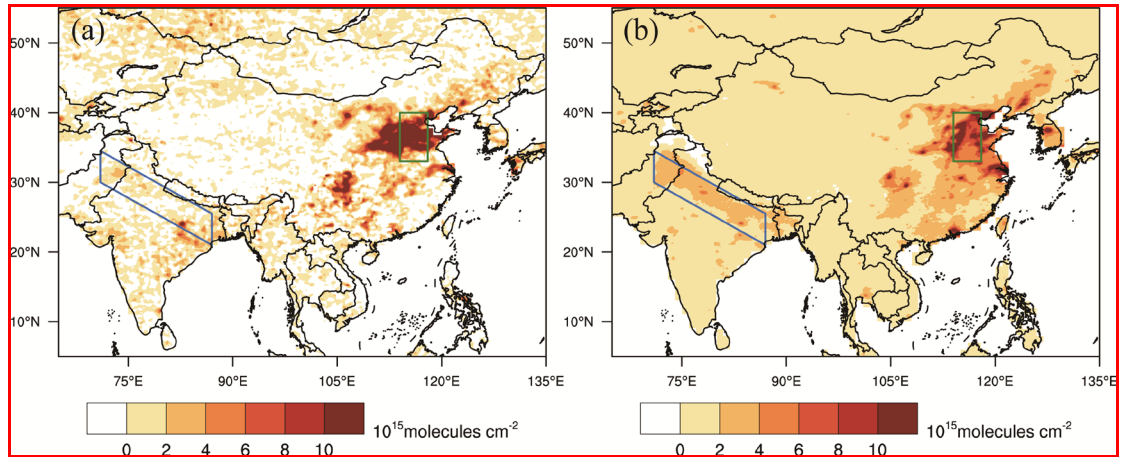
Tables S1 to S2

## Introduction

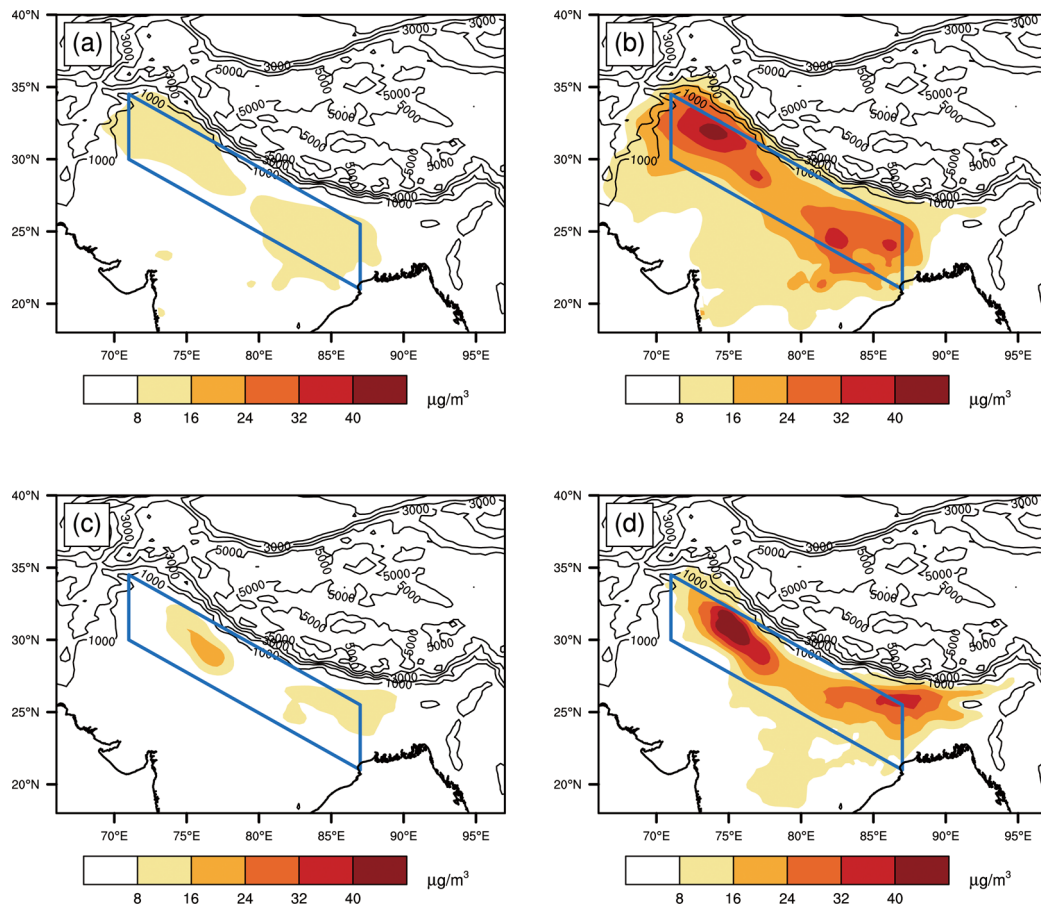
This supporting information consists of the following parts. Figure S1 shows spatial distributions of the emission fluxes of NH<sub>3</sub>, SO<sub>2</sub> and NO<sub>x</sub> over East Asia from June to August 2010 estimated using MIX database. Figure S2 shows the spatial distributions of SO<sub>2</sub> and NO<sub>2</sub> columns over East Asia. Figure S3 provides the comparison of the simulated SO<sub>4</sub><sup>2-</sup> and NO<sub>3</sub><sup>-</sup> concentrations in the base case and the increased emissions case. Figure S4 shows the spatial distribution of the NH<sub>3</sub> total columns from June to August 2010 derived from IASI measurements. Figure S5 shows spatial distributions of WRF-Chem predicted relative humidity and precipitation from June to August 2010, and the circles in Figure S5a represent the observed relative humidity obtained from NCDC dataset. Table S1 lists the options of WRF-Chem configurations. Table S2 provides the performance statistics of meteorological predictions of WRF-Chem.



**Figure S1.** Spatial distributions of emission fluxes of (a) NH<sub>3</sub>, (b) SO<sub>2</sub>, and (c) NO<sub>x</sub> over East Asia from June to August 2010. The blue quadrangle represents the IGP, and the green quadrangle represents the NCP.

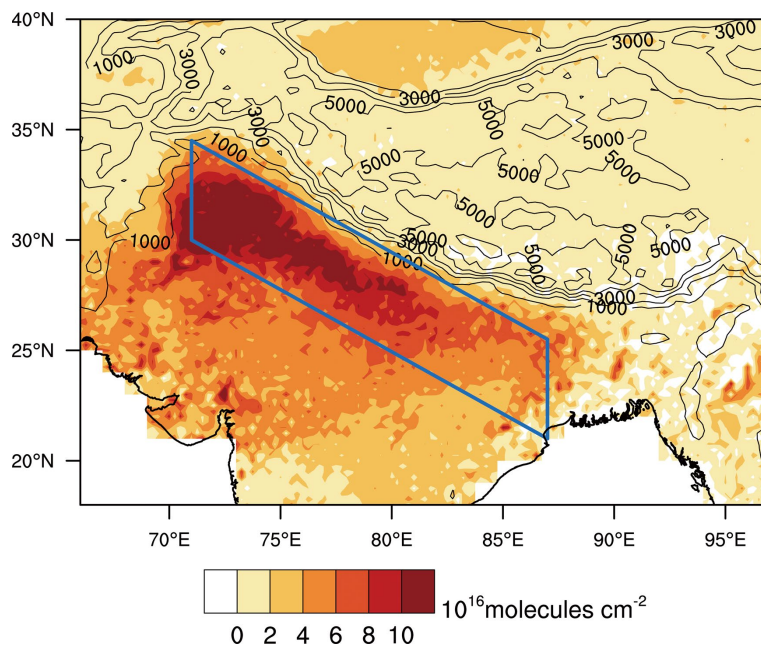


**Figure S2.** The spatial distributions of (a) SO<sub>2</sub> and (b) NO<sub>2</sub> columns over East Asia from June to August 2010.

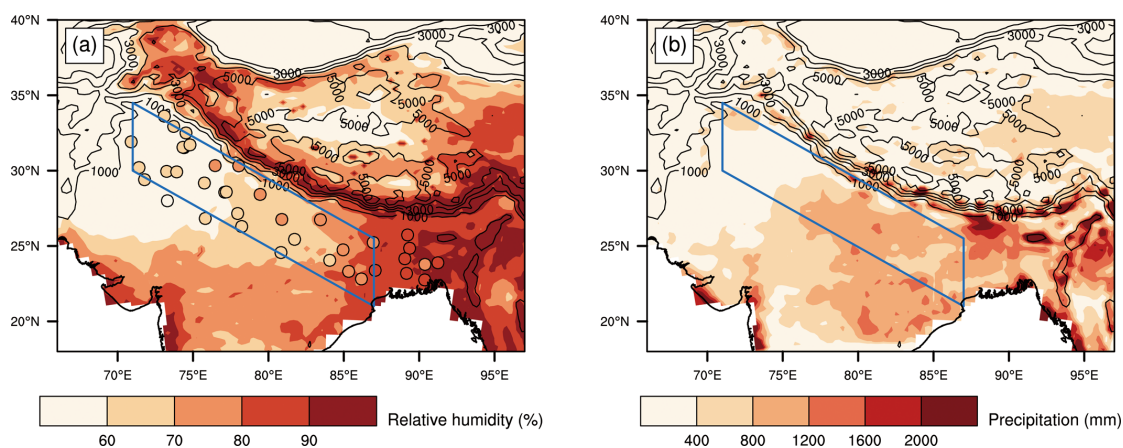


**Figure S3.** Spatial distributions of WRF-Chem predicted SO<sub>4</sub><sup>2-</sup> and NO<sub>3</sub><sup>-</sup> concentrations from June to August 2010. (a) and (b) are SO<sub>4</sub><sup>2-</sup> concentrations in the base case and the increased emissions case, respectively. (c) and (d) are NO<sub>3</sub><sup>-</sup> concentrations in the base case and the increased emissions case, respectively.





**Figure S4.** The spatial distribution of NH<sub>3</sub> total columns from June to August 2010 retrieved from IASI measurements.



**Figure S5.** Spatial distributions of WRF-Chem predicted meteorological variables from June to August 2010. (a) Relative humidity. (b) Precipitation. Circles in (a) show the observed Relative humidity.

**Table S1.** WRF-Chem configurations

Meteorology initial and boundary conditions	Reanalysis data from the National Centers for Environmental Prediction Final Analysis (NCEP-FNL)
Shortwave radiation	rapid radiative transfer model (RRTMG)
Longwave radiation	rapid radiative transfer model (RRTMG)
Land surface model	Noah land-surface model
Planetary boundary layer model	Mellor-Yamada-Janjic (Eta) TKE scheme
Cumulus parameterization	New Grell scheme (G3)
Microphysics	Lin et al. Scheme
Photolysis	Fast-J photolysis

**Table S2.** Performance statistics of meteorological predictions of WRF-Chem.

	T2 <sup>a</sup>	RH2 <sup>a</sup>	WS10 <sup>a</sup>	WD10 <sup>a</sup>
Data pairs <sup>b</sup>	27508	27443	18036	18036
MeanObs <sup>b</sup>	30.9	69.3	2.7	165.1
MeanSim <sup>b</sup>	31.9	63.3	2.9	152.9
R <sup>b</sup>	0.8	0.8	0.1	0.4
MB <sup>b</sup>	0.9	-5.9	0.2	-12.1
RMSE <sup>b</sup>	3.6	19.8	2.7	95.2
NMB (%) <sup>b</sup>	3.0	-8.6	9.1	-7.4

<sup>a</sup> T2: temperature at 2 m; RH2: relative humidity at 2 m; WS10: wind speed at 10 m; WD10: wind direction at 10 m; SLP: sea level pressure.

<sup>b</sup> data pairs: the number of observed and simulated data pairs; MeanObs: mean observational data; MeanSim: mean simulation results; R: correlation coefficient; MB: mean bias; RMSE: root mean square error; NMB: normalized mean bias.