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# 1 Relative Humidity Effect on the Formation of Highly Oxidized Molecules and

# 2 New Particles during Monoterpene Oxidation

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10 Abstract. It has been widely observed around the world that the frequency and intensity of new particle formation (NPF) 11 events are reduced during periods of high relative humidity (RH). The current study focuses on how RH affects the formation 12 of highly oxidized molecules (HOMs), which are key components of NPF and initial growth caused by oxidized organics. The 13 ozonolysis of  $\alpha$ -pinene, limonene, and  $\Delta^3$ -carene, with and without OH-scavenger, were carried out under low NOx conditions 14 under a range of RH (from ~3% to ~90%) in a temperature-controlled flow tube. A Scanning Mobility Particle Sizer (SMPS) 15 was used to measure the size distribution of generated particles and a novel transverse-ionization chemical ionization inlet with 16 a high-resolution time-of-fight mass spectrometer detected HOMs. A major finding from this work is that neither the detected 17 HOMs nor their abundance changed significantly with RH, which indicates that the detected HOMs must be formed from 18 water-independent pathways. In fact, the distinguished OH- and O<sub>3</sub>-derived peroxy radicals (RO<sub>2</sub>), HOM monomers, and 19 HOM dimers could mostly be explained by the autoxidation of RO<sub>2</sub> followed by bimolecular reactions with other RO<sub>2</sub> or 20 hydroperoxy radicals (HO<sub>2</sub>), rather than from a water-influenced pathway like through the formation of a stabilized Criegee 21 intermediate (sCI). However, as RH changed from 3 to 90% the particle number concentrations decreased by a factor of 2~3 22 while particle mass concentrations increased or decreased slightly within a factor of 2. These observations show that, while 23 high RH appears to inhibit NPF as evident by the decreasing number concentration, this reduction is not caused by a decrease 24 in RO<sub>2</sub>-derived HOMs formation. One possible explanation is the existence of other extremely low volatility compounds 25 (ELVOCs), like gas phase formed sCI-included accretion products, which are responsible for the very first steps of NPF but are 26 not detected by nitrate-based chemical ionization mass spectrometry. These ELVOCs may be preferentially reduced at high 27 RH compared to more volatile compounds, the latter of which mainly determine the final mass concentration of particles. 28 Another possibility is that a fraction of HOMs cluster with water (but detected as the declustered molecules) at high RH in such 29 a way that they may no longer be able to participate in cluster formation, thereby suppressing NPF.





#### 30 1 Introduction

31 New particle formation (NPF) is ubiquitous around the world (Kulmala et al., 2004). Newly formed particles contribute 32 greatly to global particle populations and can grow further to act as cloud condensation nuclei (CCN), thereby influencing 33 clouds and climate (Makkonen et al., 2012; Merikanto et al., 2009; Dunne et al., 2016). NPF characteristics vary from site to 34 site because of varying precursors and atmospheric conditions. It has been widely observed that the intensity (Sihto et al., 35 2006; Dada et al., 2017) and frequency (Dada et al., 2017; Boy and Kulmala, 2002; Hyvönen et al., 2005) of continental NPF 36 are reduced during periods of high RH, resulting in reduced ultrafine particle number concentrations during these periods 37 (Weber et al., 1997). For example, 20 years of observations in the boreal forest at Hyytiälä, Finland, showed that NPF is more likely during periods of low ambient RH (Dada et al., 2017). In central Amazonia, where particle composition is 38 39 dominated by oxidation products of biogenic organic compounds, new particles were not formed at ground-level where RH 40 was always higher than 60%, but rather formed in the upper troposphere where RH and condensation sink (CS) were 41 significantly lower (Poschl et al., 2010; Andreae et al., 2018). In urban areas, NPF also favors low RH (Cai et al., 2017; Shen et al., 2011). Despite the low continental NPF event frequency at high RH, NPF has still been observed in the free 42 43 troposphere in vicinity of clouds, where RH is extremely high (Weber et al., 1999) and in coastal and marine areas where RH 44 is typically greater than 90% (O'Dowd et al., 1998).

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The widely observed anti-correlation between NPF and RH in the field experiments can be attributed to the indirect 46 47 influence of water. For example, high RH often corresponds to greater cloud cover, which can lead to lower ground-level 48 concentrations of photo-oxidized precursors such as H<sub>2</sub>SO<sub>4</sub> and highly oxidized molecules (HOMs) as well as an increased 49 condensation sink that leads to scavenging of precursors and clusters (Hamed et al., 2011). On the other hand, water vapor 50 may also directly influence NPF by regulating the formation of gas phase precursors or by participating in cluster formation. 51 For example, chamber and model experiments on the binary sulfuric acid-water system have demonstrated positive 52 relationships between particle formation rate and RH (Duplissy et al., 2016; Merikanto et al., 2016). While in the ternary 53 (H<sub>2</sub>SO<sub>4</sub>/MSA-H<sub>2</sub>O-Amine/NH<sub>3</sub>) system, H<sub>2</sub>O was reported to have either positive (Chen et al., 2015) or negative (Napari et 54 al., 2002) effects on NPF. Some studies have hypothesized that high water content might suppress the formation of 55 NPF-related organics from the oxidation of biogenic precursors (Hyvönen et al., 2005; Boy and Kulmala, 2002). However, 56 no direct evidence of this has been provided.

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58 Although sulfuric acid has been recognized as the most important precursor of new particle formation, it alone can't explain





59 the rapid formation and growth rates observed in the field (Kuang et al., 2008). Organic compounds, ammonia, amines, and 60 water are also likely involved (Zhang et al., 2012; Chen et al., 2012). Organics have been shown to be very important for 61 cluster formation and stabilization in theoretical studies (Ortega et al., 2016; Donahue et al., 2013), laboratory experiments 62 (Tröstl et al., 2016; Schobesberger et al., 2013) and field measurements (Bianchi et al., 2016; Hoffmann et al., 2001; Metzger 63 et al., 2010). Organics can either form clusters with sulfuric acid or purely with themselves (Zhao et al., 2013; Zhao et al., 64 2009). They can also contribute significantly to the condensational growth of newly formed particles, determining particle 65 growth rates, particle lifetime, and global particle and CCN concentrations (Donahue et al., 2011; Vehkamäki and Riipinen, 66 2012). The ability of organics to take part in particle formation and condensational growth depends on their volatility as well 67 as reactivity. HOMs, such as extremely low volatility organic compounds (ELVOCs, saturation mass concentration ( $C^*$ ) <  $3 \times 10^{-4} \ \mu g \text{ m}^{-3}$ ) or low volatility organic compounds (LVOCs,  $3 \times 10^{-4} < \text{C}^* < 0.3 \ \mu g \text{ m}^{-3}$ ), are likely contributors to NPF 68 (Donahue et al., 2012; Ehn et al., 2014). 69

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71 Despite its large contribution to NPF, the direct measurement of HOMs has long been a challenge because of their low 72 atmospheric concentrations, low volatilities, and short lifetimes. Recently, the development of the high resolution 73 time-of-flight chemical ionization mass spectrometer (HRToF-CIMS) overcame this barrier and made the measurement and 74 identification of HOMs feasible (Junninen et al., 2010; Jokinen et al., 2012). HOMs from both monoterpene and aromatic 75 oxidation showed high O/C ratios of > 0.7-0.8, and were present as monomers, dimers and even higher order clusters 76 (Molteni et al., 2018; Ehn et al., 2012). These high O/C ratios could not be explained by any of the formerly known 77 oxidization pathways unless the autoxidation of RO2 was taken into consideration (Crounse et al., 2013; Barsanti et al., 78 2017). Autoxidation was widely observed in condensed phase reactions, however, it was not considered in the gas phase 79 previously because of the high energy barrier. This was further confirmed by the fact that at higher temperatures, more 80 HOMs are formed than at low temperatures (Frege et al., 2018).

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Since most laboratory experiments related to the formation of HOMs have been conducted under conditions of constant RH, usually low or medium RH of less than 60% (Ehn et al., 2012; Zhang et al., 2015), it was still unknown whether and how water vapor might impact HOMs formation. High RH conditions are difficult to achieve in chamber experiments without significantly changing temperature and pressure. In addition, HOMs detection by the current commercially available CIMS inlet based on the design of Eisele and Tanner is subject to water cluster influence (Kuerten et al., 2016).

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88 In this research, three different endocyclic monoterpenes,  $\alpha$ -pinene, limonene and  $\Delta^3$ -carene were reacted with ozone, with





and without hydroxyl radical (OH) scavengers, in a reaction flow tube. RH influences on HOMs formation and organic-driven NPF were studied under a range of RH from 0% to 90%. Generated closed-shell HOMs and RO<sub>2</sub> were measured using a home-built CIMS inlet coupled to a HRToF mass spectrometer (LTOF mass analyzer, Tofwerk AG). The CIMS inlet effectively reduced water clustering onto ions sampled into vacuum, thus removing sample artifacts caused by high water vapor levels. Water vapor influence on the formation of RO<sub>2</sub>, HOM monomers and HOM dimers were studied. The volatility of O<sub>3</sub>- and OH-derived closed-shell HOMs were estimated with a group contribution-based model (SIMPOL) and a recently developed statistical model to study the potential contribution of O<sub>3</sub> and OH initiated chemistry on NPF.

#### 96 2 Methods

#### 97 2.1 Flow tube reactor

98 The experiments were performed in a laminar flow tube reactor consisting of a 150 cm long Pyrex glass cylinder with a 99 volume of 8.5 dm<sup>3</sup> (Figure 1). The flow tube was located in a temperature controlled room ( $T=293\pm2^{\circ}C$ ) and was covered so 100 that all experiments were performed under dark conditions. The monoterpenes were injected into the chamber using a 101 syringe pump (model NE-300, New Era Pump Systems, Inc.) evaporated into a 2.5 LPM flow of dry, purified "zero air." O<sub>3</sub> 102 was generated by passing 0.5 LPM dry zero air (79% N2, 21% O2) over a Hg UV lamp (model 90-0004-04, UVP, LLC) and 103 then diluted with 6.5 LPM of zero air at the prescribed RH. The zero air was generated with a zero air generator (model 104 747-30, Aadco Instruments), with NOx and SO<sub>2</sub> concentrations specified to be less than 0.5 ppbv. Low NOx conditions were 105 achieved using zero air as-is. A temperature-controlled bubbler filled with deionized water was used to generate humid air, 106 and the prescribed RH was achieved by controlling the temperature of the bubbler. Gas inlets to the flow tube were made 107 from Teflon tubing that were capped and drilled with small holes to distribute sample air uniformly in the flow tube cross 108 section. The uniform distributions of O<sub>3</sub> and H<sub>2</sub>O in the flow tube were confirmed by measuring [O<sub>3</sub>] and RH at the different locations prior to the experiments. In every experiment, RH was adjusted to be constant for at least 30 min for each of the 109 110 four RH steps (0~5%, 30~38%, 58~65%, 85~92%). The total flow rate was kept at 8.5 LPM so that the average reaction time 111 was constant (~60s) for each experiment. At the beginning of the experiments, the inner wall of the reactor was washed with 112 ultra-pure water. All of the flow rates were calibrated before and during the experiments.

#### 113 2.2 Instrumentation

#### 114 2.2.1 Transverse Ionization – Chemical Ionization Mass Spectrometer

A self-designed and home-built chemical ionization inlet, called Transverse Ionization (TI) inlet (Figure 2 and Figure S1),

- 116 was used in front of the LTOF mass analyzer. The TI design is similar to those of the Ambient-pressure Proton transfer Mass
- 117 Spectrometer (AmPMS) (Hanson et al., 2011) and the cluster-CIMS (Zhao et al., 2010). In the TI inlet, a 4-10 LPM flow of





118 sample air is passed across the inlet orifice of the mass spectrometer, where it encounters an orthogonal, 1 LPM reagent ion 119 gas flow consisting N<sub>2</sub> containing ionized nitrate ions (NO<sub>3</sub><sup>-</sup>) as well as potential cluster ions (HNO<sub>3</sub>)<sub>n</sub>NO<sub>3</sub> with n=1-3. For 120 the current study, the sample flow to the inlet was set to 4.5 LPM. Chemical ionization occurs at atmospheric pressure and 121 temperature. The reagent gas is generated by passing 3 ccm of  $N_2$  over a small vial containing nitric acid, which is then ionized by a 370 MBq  $Po^{210}$  radioactive source (model P-2021, NRD, LLC). An additional flow of  $N_2$  can be added to the 122 123 reagent gas to change the reagent ion concentration, and the assembly can be adjusted to vary ion-molecule reaction time. 124 The latter can be controlled by adjusting the sample and reagent gas flow rates or by applying different voltages to the 125 ionization source and the main inlet block. To minimize the diffusion loss in sample lines, the inlet of the TI source was 126 connected to the flow tube outlet by a short (~10 cm) piece of electro-polished stainless steel tubing. Compared to the widely used commercial nitrate source patterned after the design by Eisele and Tanner (1993) and marketed by Aerodyne, Inc., no 127 128 additional sheath flow is required so thus any impurities potentially introduced by the sheath flow are eliminated. Some flow 129 disturbance may occur where the sample flow encounters the transverse reagent flow, which may lead to non-ideal behavior. 130 However, even at the maximum total flow of 11 LPM, the Reynolds number in this region is ~500 and thus turbulence is not 131 expected to be significant.

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133 Another unique aspect of the TI design is the use of an  $N_2$  curtain gas in front of the inlet orifice to the mass spectrometer to 134 reduce water clustering on reagent and sample ions. Water clusters are expected to form at high RH mainly during the 135 free-jet expansion of the sampled gas on the vacuum side of the orifice plate (Thomson and Iribarne, 1979). The presence of 136 these clusters makes the identification and quantification of both sample and reagent ions challenging (Kulmala et al., 2014; 137 Lee et al., 2014; Kuerten et al., 2016; Ehn et al., 2014). Figure 2 shows the details of the TI source that address this issue. 138 Small holes drilled in a radial channel blow N2 uniformly in front of the orifice plate so that only sampled ions and this clean  $N_2$  gas pass into the vacuum chamber. Since the sampling flow rate of the mass spectrometer is ~0.5 LPM when using 0.3 139 140 mm orifice, the N<sub>2</sub> curtain flow is set to be 1 LPM to overflow the region surrounding the orifice. By applying voltages to 141 the ion source and the block, the ions can be efficiently guided into the mass spectrometer while neutral molecules such as 142 water vapor are prevented from entering by the N2 curtain gas.

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This TI inlet is suitable to all types of reagent ion chemistry, e.g.  $NO_3^-$ ,  $\Gamma$ , and  $H_3O^+$ . Nitrate ion chemistry was used as the reagent ion in these experiments, which is selective to highly oxidized molecules that have at least two hydroperoxy (-OOH) groups or some other H-bond-donating groups (Hyttinen et al., 2015). HOM monomers, HOM dimers and highly oxidized RO<sub>2</sub> radicals can also be measured using nitrate ion chemistry.





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#### 149 2.2.2 Other measurements

Ozone concentrations were measured with two ozone analyzers (model 106L, 2B Technology) at the inlet and outlet of the 150 151 flow tube. The sampling flow of each analyzer is 1 LPM. The two ozone analyzers were intercompared prior to the 152 experiments and the difference was within 5 ppbv when  $[O_3] < 1000$  ppbv. A Scanning Mobility Particle Sizer (SMPS), consisting of a Po210 bipolar neutralizer, a nano-Differential Mobility Analyzer (nano-DMA; model 3081, TSI, Inc.), and a 153 154 condensation particle counter (MCPC; model 1720, Brechtel Manufacturing) were used to measure the number-size 155 distribution of particles, which is later used to deduce the total particle number and mass concentrations (the latter assumes a 156 uniform density for organic particles of 1.2 g cm<sup>-3</sup>). The sampling flow rate of the MCPC was 0.3 LPM and the sheath and 157 excess flows of the nano DMA were set to 3 LPM. The flow tube particle number-size distribution was measured without further drying to get a more accurate measure of the actual particle surface area and volume, which are important for HOMs 158 159 partitioning, and also to prevent particle evaporation during the measurements.

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#### 161 2.3 Experimental conditions

Three monoterpenes were used in our experiments (see Table 1),  $\alpha$ -pinene, limonene and  $\Delta^3$ -carene. Oxidation by ozone is 162 163 believed to dominate over other oxidation radicals (i.e., OH or NO<sub>3</sub>) in forming secondary organic aerosol (SOA) under 164 atmospheric conditions (Atkinson and Arey, 2003). Ozonolysis of alkenes generates a substantial amount of OH, leading to 165 products that are produced by a combination of  $O_3$  and OH oxidation. For some experiments, in order to isolate oxidation by 166 O<sub>3</sub>, cyclohexane (see Table 1 for mixing ratios) was premixed with the monoterpene and added to the flow tube as an OH 167 scavenger. For other experiments, the combination of OH and O<sub>3</sub> chemistry were investigated to study atmospheric oxidation 168 chemistry more representative of ambient air. The "high concentration" experiments were conducted with similar mixing ratios of monoterpene ( $\sim$ 1100ppb) and O<sub>3</sub> ( $\sim$ 900ppb). The "low concentration" experiments were conducted to study the 169 170 particle-free chemical processes with initial concentrations of monoterpenes and O<sub>3</sub> shown in Table 1. Since wall losses 171 should be comparable for different precursors as a function of RH, it was not taken into consideration in our analysis of 172 HOMs production.

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#### 174 2.4 HOMs volatility predictions

The SIMPOL.1 method (Pankow and Asher, 2008) and the molecular corridor method (Li et al., 2016) were used to predict the saturation mass concentrations ( $C^*$ ) of some of the detected OH- and O<sub>3</sub>-related HOMs. SIMPOL.1 is a group contribution method and requires information on molecular structure, while the molecular corridor method only requires the





178 molecular formulae. Both methods are semi-empirical and based on volatility data from hundreds or thousands of 179 compounds. The calculated volatilities were then applied to the two-dimensional volatility basis set (2D-VBS) (Donahue et 180 al., 2012) to explore the likelihood that the products participate in the initial stages of nanoparticle growth.

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### 182 3 Results and discussion

#### 183 3.1 TI-CIMS performance

184 When comparing the TI inlet with the commercial nitrate inlet in measuring  $\alpha$ -pinene ozonolysis products, both inlets 185 produced identical mass spectra. The sensitivities of both inlets to H<sub>2</sub>SO<sub>4</sub> were determined using a home-built H<sub>2</sub>SO<sub>4</sub> 186 calibration system (Figure S2) based on the design of Kurten et al. (2012). Figure 3 summarizes the results of these 187 calibrations. The position of the ion source relative to the inlet orifice is critical for determining the sensitivity of the TI inlet. 188 When the ion source is placed 0.5 cm upstream along the sample flow axis and 5 cm away from the inlet orifice along the 189 reagent ion flow axis (configuration shown in Figure 2), the instrument is at its most sensitive. The calibration factors, defined as  $C = [H_2SO_4]/([HSO_4^-])/[NO_3^-])$  (Eisele and Tanner, 1993), for the TI in this position and the commercial inlet 190 were  $3.25 \times 10^{10}$  molecules cm<sup>-3</sup> and  $1.41 \times 10^{10}$  molecules cm<sup>-3</sup>, respectively. The lower calibration factor for the TI is 191 192 attributed to the shorter reaction time, which we estimate to be ~80 ms. The total ion count (TIC) of the TI inlet is more than 193 5 times higher than the commercial inlet, so the overall sensitivity is better. The lower detection limit for sulfuric acid, which 194 is defined as three times the standard deviation of the background (Jokinen et al., 2012), is  $9.3 \times 10^4$  molecules cm<sup>-3</sup> and  $1.26 \times 10^5$  molecules cm<sup>-3</sup> for the TI and commercial inlet, respectively. 195

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197 After applying the N2 curtain gas flow, the TIC recorded by the TI-CIMS decreased significantly. This was compensated for by increasing the ion source and reaction chamber voltages that direct ions to the orifice (Figure S3). When RH>90%, the 198 199 reagent ion mass spectrum was dominated by water clusters  $(H_2O)_m(HNO_3)_nNO_3^-$  (m=0~30, n=0~2) if no N<sub>2</sub> curtain flow 200 was applied. The reagent ions NO3<sup>-</sup>, HNO3NO3<sup>-</sup> and (HNO3)2NO3<sup>-</sup> decreased as RH increased, with [(HNO3)2NO3<sup>-</sup>] and 201 [HNO<sub>3</sub>NO<sub>3</sub><sup>-</sup>] decreasing much faster than [NO<sub>3</sub><sup>-</sup>]. In contrast, after 1 LPM N<sub>2</sub> curtain flow was applied to the inlet, most of 202 the water clusters were removed (Figure 4). The reagent ions, sample ions and TIC remained stable as RH increased, which 203 resulted in a reliable measurement of HOMs concentrations as a function of RH. The result that the N2 curtain flow 204 eliminated water clustering to a large extent confirms that most of the water clusters in the spectrum were produced during 205 the free-jet expansion into vacuum instead of formed in the ion-molecular reagent region.





#### 207 3.2 Identification of HOMs spectrum

208 Figure 5 shows the average mass spectra of the HOMs dimers and Figure S4 shows the average mass spectra of the HOMs 209 monomer and RO<sub>2</sub> radicals for each of the six particle generation experiments. More than 400 peaks were identified in each spectrum, the majority of which were clusters with NO<sub>3</sub><sup>-</sup> or HNO<sub>3</sub>NO<sub>3</sub><sup>-</sup>. [H<sub>2</sub>SO<sub>4</sub>] was  $\sim 10^5$  molecules cm<sup>-3</sup> and was always 210 211 less than 3% of the most abundant  $C_{10}$  products, suggesting that sulfuric acid plays a negligible role in nucleation and cluster 212 growth in our experiments. After subtracting the reagent ions (NO3<sup>-</sup> or HNO3NO3<sup>-</sup>), molecular formulae for organics with an 213 odd number of H atoms were assigned to radicals, which are generally difficult to detect experimentally (Rissanen et al., 214 2015), and formulae with an even number of H atoms were assigned to closed-shell molecules. Most of the HOMs products 215 from the three endocyclic monoterpenes were very similar, while the relative abundance of different HOMs was quite different, indicating similar reaction pathways but different branching ratios in the reaction mechanisms. The main products 216 were C<sub>5-10</sub>H<sub>6-16</sub>O<sub>3-10</sub> for closed shell monomers and RO<sub>2</sub> and C<sub>15-20</sub>H<sub>22-34</sub>O<sub>6-18</sub> for closed shell dimers. Among these, C<sub>10</sub> and 217 218 C20 compounds were the most abundant. C5.9 products could be formed from O3 attack on the less reactive exocyclic carbon 219 double bond or the decomposition of intermediate radicals. Some fragments were found to be unique for specific 220 monoterpene precursors. For instance,  $C_5H_6O_7$  (m/z 240) was much more abundant in  $\alpha$ -pinene oxidation than from other 221 two precursors, which might be a tri-carboxylic acid (Ehn et al., 2012).

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223 Comparing total HOMs abundance for the three monoterpene oxidation reactions, limonene created the most, followed by 224  $\alpha$ -pinene and then  $\Delta^3$ -carene. This is in qualitative agreement with prior studies (Jokinen et al., 2014; Ehn et al., 2014). The 225 total dimer signal intensity was 15-30% of monomers for all three monoterpenes. Experiments with an OH scavenger 226 generated fewer HOMs than those without OH scavengers.

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228 As observed in previous studies, C<sub>10</sub>H<sub>15</sub>O<sub>6,8,10,12</sub> and C<sub>10</sub>H<sub>17</sub>O<sub>5,7,9,11</sub> comprised the O<sub>3</sub>- and OH-related RO<sub>2</sub>, respectively 229 (Jokinen et al., 2014), while  $C_{10}H_{14}O_{5,7,9,11}$  and  $C_{10}H_{16}O_{6,8,10,12}$  comprised the O<sub>3</sub>- and OH-related closed shell monomers, 230 respectively (Ehn et al., 2014). When comparing the average spectra with and without OH scavenger, no obvious differences were seen for OH-related RO<sub>2</sub> or monomers (Figure S4). In contrast, for dimers we found that  $C_{20}H_{32}O_{6-13}$  were more 231 232 abundant in experiments without OH scavenger (Figure 5). The formation of these dimers can be explained by the reaction 233 of one OH-related RO<sub>2</sub> with one O<sub>3</sub>-derived RO<sub>2</sub> (see Section 3.5), and can therefore be considered as markers for combined 234 OH and O<sub>3</sub> chemistry. As HOMs dimers are generally less volatile than monomers with identical O/C ratio, rapid production 235 of dimers is believed to play a more important role in initial particle formation and growth (Zhang et al., 2015).





#### 237 3.3 RH influence on HOMs generation

Figure 6 shows a time series of experimental parameters, particle size distribution, and key ions from the limonene ozonolysis experiment with OH scavenger (EXP. 2 in Table 1). The  $O_3$  inlet and outlet concentrations were approximately constant with increasing RH (Figure 6a), indicating that RH did not significantly change  $O_3$  levels in the flow tube. This also shows that the reactivity of the limonene with ozone does not change with RH. The number concentration of the generated particles decreased from  $4.9 \times 10^6$  cm<sup>-3</sup> to  $2.7 \times 10^6$  cm<sup>-3</sup> with increasing RH, while the peak of the number-size distribution increased slightly, due in part to water absorption. When RH was above 80%, both the integrated number and mass concentrations, which were calculated from the number-size distributions, decreased (Figure 6b).

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Despite the change in particle number and mass concentrations with RH, the concentration of all the main HOMs, including 246 RO<sub>2</sub>, monomers and dimers, did not change for both OH- and O<sub>3</sub>-derived products (Figure 6c). In fact, the only signals in the 247 248 mass spectra that changed with RH corresponded to an increases associated with water clusters. The variations in HOMs 249 concentrations can be explained by the competition between production and condensational losses. With the changes of particle concentration, the condensation sink should also change with RH. However, in this experiment, as in all other 250 experiments, the surface area for existing particles  $(1.8 \sim 4.1 \times 10^{-6} \text{ m}^2)$  was much lower than the wall surface area of the flow 251 252 tube (~0.38 m<sup>2</sup>). As almost all of the detected HOMs are ELVOCs or LVOCs (see Section 3.6), they are not likely to 253 partition back to the gas phase after they encounter a wall. As a result, the main loss in the flow tube should be caused by 254 wall loss, which does not change significantly with RH. To further test this hypothesis that wall losses dominated over 255 condensation onto particles, particle free experiments were performed and, again, the detected HOMs signals did not change 256 with RH (Figure 7).

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#### 258 3.4 RH influence on SOA generation

259 Figure 8 shows the integrated SOA particle number and mass concentrations. The generated SOA particle number and mass 260 concentrations for limonene were ~5 times greater than for  $\Delta^3$ -carene and  $\alpha$ -pinene. This is because the theoretical ozone 261 reactivity of limonene is 3~5 times higher than the latter two and molar yield from limonene ozonolysis is also the highest. 262 Peaks in the particle number-size distributions were between 40 and 70 nm (Figure S5). In most of the experiments, generated SOA mass concentrations increased or decreased slightly when RH increased from ~0% to ~60% and decreased as 263 264 RH further increased to ~90%. The variability in particle mass concentration as a function of RH for different experiments 265 can be attributed to combined effects of gas phase reactions, condensed phase reactions, as well as physical uptake of water, 266 but all values were within a factor of two of each other. On the contrary, the particle number concentrations decreased by a





267 factor of 2~3 with increasing RH.

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269 A number of studies have demonstrated different water and OH influences on the ozonolysis products of exocyclic and 270 endocyclic organic compounds. They have reported either suppressing (Bonn et al., 2002; Bonn and Moorgat, 2002) or 271 promoting (Jonsson et al., 2006; Jonsson et al., 2008) effects of water vapor on the particle formation processes during 272 ozonolysis of monoterpenes by measuring the number-size distributions of generated SOA particles with SMPS. The 273 discrepancies between different results could be attributed to the different experiment setups, e.g., monoterpene and O<sub>3</sub> 274 concentration, temperature, RH range, OH scavengers, reaction time, and so on. Specifically, our results are in good 275 agreement with those of Bonn et al., who studied SOA generation from the ozonolysis of endocyclic monoterpenes (Bonn et al., 2002). In that study, SOA number concentrations decreased by a factor of 1.1~2.5 as RH increased, while the variation in 276 277 volume concentrations was negligible (within  $\pm 10\%$ ). They concluded that water's influence on non-volatile products, which 278 are responsible for the initial steps of nucleation, was much larger than its influence for semi-volatile compounds which 279 mainly determined the final volume concentrations of particles. Thus, it was highly suspected that water influenced new 280 particle formation through influencing the generation of NPF precursors. However, our measurements indicate that at least 281 the formation of the detected HOMs is independent of water vapor concentrations. There may be other species that are 282 crucial to the initial steps of NPF and are affected by water vapor but are not detected by nitrate CIMS (see section 3.5). 283 Another possible explanation is that a fraction of HOMs cluster with water at high RH in such a way that they may no longer 284 be able to participate in further cluster formation, thereby suppressing NPF. If the CIMS measurement only detected the 285 declustered molecule, then such a mechanism may still be consistent with our observations.

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# 287 3.5 Possible water-relevant $C_{10}$ and $C_{20}$ HOMs formation pathways

Although the oxidation of BVOCs has been widely studied, it has mostly been constrained to the early stages (first and 288 289 second generation intermediates) and many uncertainties still exist (Johnson and Marston, 2008; Isaacman-VanWertz et al., 290 2018; Atkinson and Arey, 2003). The first step of ozonolysis for the three BVOCs ( $\alpha$ -pinene,  $\Delta^3$ -carene and limonene) is 291 ozone attack on the endocyclic carbon double bond to form a primary ozonide. Figure 9 shows the O3-initiated oxidation 292 pathways of α-pinene that may be related to the detected C<sub>10</sub> and C<sub>20</sub> HOMs for representative isomers. The primary ozonide 293 rapidly transforms to two excited Criegee intermediates (eCIs), one of which (branching ratio= 0.4) (Kamens et al., 1999) is 294 shown in Figure 9. The reaction pathways of the eCI are complex, the most important two under ambient and most chamber 295 conditions are the sCI channel (reaction I) and the hydroperoxide channel (reaction II) (Bonn et al., 2002). The sCI either 296 reacts with aldehydes to form a secondary ozonide (when the aldehyde is  $C_{10}$ , then the formed SOZ is  $C_{20}$  and is marked as





297 sCI-C<sub>10</sub>) or with water or other acidic compounds such as alcohols and carboxylic acids to form hydroxy-hydroperoxide, 298 which then decomposes to carboxylic acids or aldehydes. For a-pinene, the main decomposition product is pinonic acid. In 299 the hydroperoxide channel (reaction II), the formed hydroperoxide quickly decomposes to a first generation alkyl radical (R) 300 and OH (Johnson and Marston, 2008). R reacts with  $O_2$  immediately to form the first generation  $RO_2$ , which can undergo 301 numerous reactions, including reaction with HO<sub>2</sub>, R'O<sub>2</sub> and autoxidation. The reaction with HO<sub>2</sub> mainly forms 302 hydroperoxides, with a small fraction forming peroxides or carbonyl-containing compounds. However, the carbonyl cannot 303 be formed in most cases since the C atom bonded to O-O does not have available electrons for the carbonyl  $\pi$ -bond. When 304 reacted with another R'O2, either ROOR' or an alkoxy radical (RO) or a carbonyl and a hydroperoxide are formed. The RO 305 can undergo isomerization, or form a carbonyl and HO<sub>2</sub>, for which the branching ratios are extremely difficult to evaluate. 306 RO can also undergo decomposition, which is one of the pathways to form  $C_5 \sim C_9$ . The autoxidation process is key to HOMs 307 formation. Each autoxidation step adds two O atoms to the molecule and thus increases the oxidation state very rapidly. The 308 competition between autoxidation processes and bimolecular reactions (RO2 reactions with R'O2 or HO2) determines the 309 ultimate oxidation state of the products (Barsanti et al., 2017; Crounse et al., 2013; Rissanen et al., 2015).

310

OH can be generated in the ozonolysis of alkenes and the yield is near unity (Atkinson, 1997). The reaction of OH with  $\alpha$ -pinene directly forms first generation R and then RO<sub>2</sub>; one possible structure for this RO<sub>2</sub> (branching ratio = 0.44), formed from OH addition to the double bond (Berndt et al., 2016), is shown in Figure 9. However, the formed RO<sub>2</sub> (C<sub>10</sub>H<sub>17</sub>O<sub>2m+1</sub>) are not the same RO<sub>2</sub> as those formed through ozonolysis (C<sub>10</sub>H<sub>15</sub>O<sub>2n+2</sub>) (McVay et al., 2016). Accordingly, the structure and composition of C<sub>10</sub> and C<sub>20</sub> HOMs formed from OH or O<sub>3</sub> chemistry are different, and so too are their potential impacts on NPF. The combined OH- and O<sub>3</sub>-derived dimers (C<sub>10</sub>H<sub>32</sub>O<sub>2(m+n)+3</sub>), formed by collision of an OH-derived RO<sub>2</sub> with an O<sub>3</sub>-derived RO<sub>2</sub>, were only observed in ozonolysis experiments without OH scavenger.

318

The RO<sub>2</sub> autoxidation pathway explains most of the observed C<sub>10</sub> and C<sub>20</sub> compounds in the mass spectra. One exception to this is that C<sub>10</sub>H<sub>18</sub>O<sub>2m+1</sub>, C<sub>10</sub>H<sub>18</sub>O<sub>2m</sub> and C<sub>20</sub>H<sub>34</sub>O<sub>2(m+m<sup>+</sup>)</sub> were not observed in the spectrum whereas in experiments performed by Berndt et al., in which OH oxidation for  $\alpha$ -pinene was studied, C<sub>10</sub>H<sub>18</sub>O<sub>2m</sub> and C<sub>20</sub>H<sub>34</sub>O<sub>2(m+m<sup>+</sup>)</sub> dominated the mass spectrum (Berndt et al., 2016). This could be explained by a low OH/O<sub>3</sub> ratio in our experiments, since unlike Berndt et al. we did not provide a source of OH to the flow tube.

324

325 Despite this close agreement achieved by the  $RO_2$  autoxidation mechanism and the observed mass spectra in our study, prior 326 studies suggest that other potential pathways cannot be excluded. An accretion product involving sCI is one possibility





327 (Barsanti et al., 2017). It is possible that sCI reacts with long-chain carboxylic acids or carbonyls, such as those with 10 328 carbon atoms, forming in this instance anhydrides (sCI-C<sub>10</sub>, reaction IV) or secondary ozonides (sCI-C<sub>10</sub>, reaction V) with 329 vapor pressures lower than  $10^{-15}$  torr (Kamens et al., 1999; Tobias and Ziemann, 2001; Bonn et al., 2002). The formation of 330 anhydride is more likely in condensed phase, whereas there is also a possibility it can also happen in gas phase (Kamens et 331 al., 1999). However, it is unknown whether these sCI-C<sub>10</sub> can be detected using nitrate-CIMS as they may lack hydrogen 332 bond donor moieties. The semi-volatile pinonic acid can also form HOMs after further oxidation by OH (Ehn et al., 2014), 333 provided that excess α-pinene is not present to compete with pinonic acid for the generated OH.

334

335 Water vapor's influence on HOMs formation can be direct or indirect. For monoterpene oxidation, the direct participation of water vapor is to react with sCI, favoring the formation of the hydroperoxide and its decomposition products (reaction III) 336 337 over the secondary ozonides (sCI-C<sub>10</sub>, reaction IV) or possible anhydrides (sCI-C<sub>10</sub>, reaction V). Since the formation of 338 sCI-C<sub>10</sub> is more likely to contribute to NPF than the products from sCI and water vapor (Kamens et al., 1999; Tobias et al., 339 2000), a decrease in low volatility sCI-C<sub>10</sub> with high RH could explain the decreasing SOA number concentrations in our 340 experiment. It has been shown previously that OH yields from the reactions of  $O_3$  with a series of monoterpenes were not 341 affected by the presence of water vapor (Atkinson et al., 1992; Aschmann et al., 2002), which implies that the hydroperoxide 342 channel (reaction II) are similarly unaffected by water. Since the detected HOMs in our experiments were RH-independent, 343 we conclude that all the detected HOMs were formed from hydroperoxide channel (reaction II) and not via the sCI channel 344 (reaction I). Similarly, the detected HOMs were not likely to form through the hydration reaction (Equation 1) (Ehn et al., 345 2012), which is supposed to increase with increasing RH.

346

 $R - CHO + H_2O \rightarrow R - CH(OH)_2$  . (Equation 1)

347

The indirect water effect on HOMs formation includes the water influence on HO<sub>2</sub> fate. As water promotes HO<sub>2</sub> self-reaction (Equation 2), reaction of HO<sub>2</sub> with RO<sub>2</sub> should decrease and the related HOM monomers should likewise decrease with increasing RH. However, as the formation of both HOM monomers and dimers was not affected by H<sub>2</sub>O, it was likely that water does not significantly increase HO<sub>2</sub> self-reaction or that HO<sub>2</sub> chemistry was not important in our experiments.

$$2HO_2 \xrightarrow{M, H_2O} H_2O_2 + O_2 \qquad (Equation 2)$$

353

#### 354 3.6 Volatility predictions

The volatility of the gas phase products is one of the most important properties that determines whether a compound contributes to the formation, initial growth or further growth of SOA particles (Donahue et al., 2012; Kroll et al., 2011). As





357 the products with identical elemental composition can be formed from different bimolecular reactions of the intermediate 358 RO<sub>2</sub>, it is difficult to predict their exact structures. For the current study, the number of different structural and functional 359 groups (e.g., aromatic rings, carbon double bonds, aldehydes, carbonyls, hydroxyls, ethers, hydroperoxyls) was estimated 360 and used to derive saturation vapor pressure using SIMPOL.1 (Table S1). To simplify the calculation, the functional groups 361 used here were directly predicted from the proposed formation pathways in Figure 10 and did not include intramolecular 362 isomerization, although that may be important in some situations. For example, one of the ROOH can be replaced with an 363 endo-peroxide via ring closure of unsaturated RO<sub>2</sub> (Berndt et al., 2016). To figure out the possible bias introduced by this 364 simplification, the result was compared to those obtained using the Molecular Corridor method (Li et al., 2016), the latter of 365 which does not require information on functional groups.

366

Figure 10 shows the predicted saturation mass concentrations,  $C^*$ , of the main  $C_{10}$  and  $C_{20}$  closed shell products. The 367 368 difference of C\* predicted from the two methods was within one order of magnitude for C20 HOMs, and 3~4 orders of 369 magnitude for C<sub>10</sub> HOMs. The difference was due to the abundance of -OOH and -OH moieties, which contribute more to 370 lowering saturation vapor pressures than other functional groups (Table S1). Despite these differences, nearly all of the C20 371 HOMs can be classified as ELVOCs, while  $C_{10}$  products were mostly LVOCs. Typically, the volatilities of O<sub>3</sub>-derived  $C_{20}$ 372 HOMs were less than OH-related HOMs, whereas for those with identical oxidation states of carbon,  $\overline{OSc}$  (defined as 373 2O/C-H/C) (Kroll et al., 2011), such as C<sub>20</sub>H<sub>30</sub>O<sub>10</sub> (O<sub>3</sub>-derived dimer), C<sub>20</sub>H<sub>32</sub>O<sub>11</sub> (OH and O<sub>3</sub> combined dimer), C<sub>20</sub>H<sub>34</sub>O<sub>12</sub> 374 (OH-derived dimer), OH-derived C<sub>20</sub> HOMs have lower volatilities than O<sub>3</sub>-derived HOMs due to a greater number of 375 oxygen atoms in the former.

#### 376 4 Conclusions

377 The RH influence on HOMs formation and NPF during monoterpene oxidation was explored in this study. HOMs were 378 detected with a TI-CIMS, using nitrate as reagent ions; C<sub>10</sub> and C<sub>20</sub> dominated the spectra. There are mainly three potential 379 paths for water vapor influence on the formation of  $C_{10}$  and  $C_{20}$  HOMs. One is water reacting with sCI (Equation 1), thereby 380 influencing the branching ratio between formation of more volatile compounds decomposed from hydroxyl hydroperoxide, 381 such as pinonic acid, and accretion products with sCI such as secondary ozonide (sCI- $C_{10}$ ) and anhydride (sCI- $C_{10}$ ). The 382 second hypothesized water influence is on the HOMs formed from hydration reactions (Equation 2). The third is that water 383 increases the rate of self-reaction of HO<sub>2</sub> (Equation 3), thus indirectly impacts the loss pathways of RO<sub>2</sub>. Our experimental 384 results, both with high particle loading and particle-free conditions, demonstrated that neither the detected HOMs species nor their signal abundance changed significantly with RH. This indicates that the detected HOMs, which can mostly be 385





386 explained by RO2 autoxidation, must be formed from water-independent pathways rather than by those reactions mentioned 387 above. One implication of this result is that  $HO_2$  self-reaction was not significantly promoted by water or that the  $RO_2$ reaction with HO2 was not be significant in our system, but instead that RO2 reacts with another peroxy radical, R'O2, to 388 389 generate both closed shell monomers and dimers. Another implication is that the sCI pathway is not responsible for the 390 generation of the detected HOMs while the role of sCI-related HOMs (SOZ or anhydride) formation by accretion with long 391 chain products, which may not be detected with nitrate CIMS, may be important in causing the decrease in SOA number 392 concentrations with increased RH. Another possible explanation for the decreasing SOA number concentration is that water 393 may cluster with HOMs and suppress NPF.

394

The detected HOMs, which could mostly be explained by autoxidation of  $RO_2$  followed by reactions with  $R'O_2$  or  $HO_2$ , were distinguished as OH-related, O<sub>3</sub>-related RO<sub>2</sub>, closed shell HOM monomers, and HOM dimers. The volatility of the identified products were estimated with the SIMPOL.1 group contribution method and with the molecular corridor technique. That analysis confirmed that  $C_{20}$  closed shell products have significantly lower volatility compared to  $C_{10}$  products and are thus more likely to contribute to NPF. Pure O<sub>3</sub> chemistry produced lower volatility  $C_{20}$  closed shell products compared to processes that were influenced or dominated by OH. As a result, O<sub>3</sub> chemistry is suspected to be more likely to lead to NPF than OH chemistry, given the same level of oxidants and VOCs precursors.

402

#### 403 Acknowledgements

This research was supported by the US Department of Energy's Atmospheric System Research program under grant no. DESC0014469, the US National Science Foundation under grant no. AGS-1762098, and by the National Key R&D Program of China under grant no. 2017YFC0209503. XL thanks the financial support from the State Scholarship Fund managed by Chinese Scholarship Council (CSC). We thank Hayley Glicker, Deanna Caroline Myers, Michael Lawler, and Danielle Draper for their kind help.

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### 410 References

- 411 Andreae, M. O., Afchine, A., Albrecht, R., Holanda, B. A., Artaxo, P., Barbosa, H. M., Borrmann, S., Cecchini, M. A., Costa,
- 412 A., and Dollner, M.: Aerosol characteristics and particle production in the upper troposphere over the Amazon Basin, Atmos
- 413 Chem Phys, 18, 921-961, 2018.
- 414 Aschmann, S. M., Arey, J., and Atkinson, R.: OH radical formation from the gas-phase reactions of O-3 with a series of





- 415 terpenes, Atmos Environ, 36, 4347-4355, 10.1016/s1352-2310(02)00355-2, 2002.
- 416 Atkinson, R., Aschmann, S. M., Arey, J., and Shorees, B.: Formation of OH radicals in the gas phase reactions of O3 with a
- 417 series of terpenes, Journal of Geophysical Research: Atmospheres, 97, 6065-6073, 1992.
- 418 Atkinson, R.: Gas-phase tropospheric chemistry of volatile organic compounds: 1. Alkanes and alkenes, Journal of Physical
- 419 and Chemical Reference Data, 26, 215-290, 1997.
- 420 Atkinson, R., and Arey, J.: Gas-phase tropospheric chemistry of biogenic volatile organic compounds: a review, Atmos
- 421 Environ, 37, 197-219, 2003.
- 422 Barsanti, K. C., Kroll, J. H., and Thornton, J. A.: Formation of Low-Volatility Organic Compounds in the Atmosphere:
- 423 Recent Advancements and Insights, Journal of Physical Chemistry Letters, 8, 1503-1511, 10.1021/acs.jpclett.7b02969, 2017.
- 424 Berndt, T., Richters, S., Jokinen, T., Hyttinen, N., Kurtén, T., Otkjær, R. V., Kjaergaard, H. G., Stratmann, F., Herrmann, H.,
- and Sipilä, M.: Hydroxyl radical-induced formation of highly oxidized organic compounds, Nature communications, 7,
  13677, 2016.
- 427 Bianchi, F., Tröstl, J., Junninen, H., Frege, C., Henne, S., Hoyle, C. R., Molteni, U., Herrmann, E., Adamov, A., and
- 428 Bukowiecki, N.: New particle formation in the free troposphere: A question of chemistry and timing, Science, aad5456, 429 2016.
- 430 Bonn, B., and Moorgat, G.: New particle formation during a-and b-pinene oxidation by O 3, OH and NO 3, and the influence
- 431 of water vapour: particle size distribution studies, Atmos Chem Phys, 2, 183-196, 2002.
- 432 Bonn, B., Schuster, G., and Moortgat, G. K.: Influence of water vapor on the process of new particle formation during
- 433 monoterpene ozonolysis, The Journal of Physical Chemistry A, 106, 2869-2881, 2002.
- Boy, M., and Kulmala, M.: Nucleation events in the continental boundary layer: Influence of physical and meteorological
- 435 parameters, Atmos Chem Phys, 2, 1-16, 2002.
- Cai, R., Yang, D., Fu, Y., Wang, X., Li, X., Ma, Y., Hao, J., Zheng, J., and Jiang, J.: Aerosol surface area concentration: a
  governing factor in new particle formation in Beijing, Atmos Chem Phys, 17, 12327, 2017.
- 438 Chen, H., Ezell, M. J., Arquero, K. D., Varner, M. E., Dawson, M. L., Gerber, R. B., and Finlayson-Pitts, B. J.: New particle
- formation and growth from methanesulfonic acid, trimethylamine and water, Physical Chemistry Chemical Physics, 17,
  13699-13709, 2015.
- 441 Chen, M., Titcombe, M., Jiang, J., Jen, C., Kuang, C., Fischer, M. L., Eisele, F. L., Siepmann, J. I., Hanson, D. R., and Zhao,
- J.: Acid-base chemical reaction model for nucleation rates in the polluted atmospheric boundary layer, Proceedings of the
   National Academy of Sciences, 109, 18713-18718, 2012.
- 444 Crounse, J. D., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., and Wennberg, P. O.: Autoxidation of organic compounds in
- the atmosphere, The Journal of Physical Chemistry Letters, 4, 3513-3520, 2013.
- 446 Dada, L., Paasonen, P., Nieminen, T., Buenrostro Mazon, S., Kontkanen, J., Peräkylä, O., Lehtipalo, K., Hussein, T., Petäjä,
- 447 T., and Kerminen, V.-M.: Long-term analysis of clear-sky new particle formation events and nonevents in Hyytiälä, Atmos
- 448 Chem Phys, 17, 6227-6241, 2017.
- Donahue, N. M., Trump, E. R., Pierce, J. R., and Riipinen, I.: Theoretical constraints on pure vapor-pressure driven
   condensation of organics to ultrafine particles, Geophys Res Lett, 38, 10.1029/2011gl048115, 2011.
- 451 Donahue, N. M., Kroll, J., Pandis, S. N., and Robinson, A. L.: A two-dimensional volatility basis set-Part 2: Diagnostics of
- 452 organic-aerosol evolution, Atmos Chem Phys, 12, 615-634, 2012.
- 453 Donahue, N. M., Ortega, I. K., Chuang, W., Riipinen, I., Riccobono, F., Schobesberger, S., Dommen, J., Baltensperger, U.,
- 454 Kulmala, M., Worsnop, D. R., and Vehkamaki, H.: How do organic vapors contribute to new-particle formation?, Faraday
- 455 discussions, 165, 91-104, 10.1039/c3fd00046j, 2013.
- 456 Dunne, E. M., Gordon, H., Kuerten, A., Almeida, J., Duplissy, J., Williamson, C., Ortega, I. K., Pringle, K. J., Adamov, A.,
- 457 Baltensperger, U., Barmet, P., Benduhn, F., Bianchi, F., Breitenlechner, M., Clarke, A., Curtius, J., Dommen, J., Donahue, N.
- 458 M., Ehrhart, S., Flagan, R. C., Franchin, A., Guida, R., Hakala, J., Hansel, A., Heinritzi, M., Jokinen, T., Kangasluoma, J.,
- 459 Kirkby, J., Kulmala, M., Kupc, A., Lawler, M. J., Lehtipalo, K., Makhmutov, V., Mann, G., Mathot, S., Merikanto, J.,

Atmospheric Chemistry and Physics Discussions



- 460 Miettinen, P., Nenes, A., Onnela, A., Rap, A., Reddington, C. L. S., Riccobono, F., Richards, N. A. D., Rissanen, M. P.,
- 461 Rondo, L., Sarnela, N., Schobesberger, S., Sengupta, K., Simon, M., Sipila, M., Smith, J. N., Stozkhov, Y., Tome, A., Trostl,
- 462 J., Wagner, P. E., Wimmer, D., Winkler, P. M., Worsnop, D. R., and Carslaw, K. S.: Global atmospheric particle formation
- 463 from CERN CLOUD measurements, Science, 354, 1119-1124, 10.1126/science.aaf2649, 2016.
- 464 Duplissy, J., Merikanto, J., Franchin, A., Tsagkogeorgas, G., Kangasluoma, J., Wimmer, D., Vuollekoski, H., Schobesberger,
- 465 S., Lehtipalo, K., and Flagan, R.: Effect of ions on sulfuric acid water binary particle formation: 2. Experimental data and
- 466 comparison with QC normalized classical nucleation theory, Journal of Geophysical Research: Atmospheres, 121,
- 467 1752-1775, 2016.
- 468 Ehn, M., Kleist, E., Junninen, H., Petäjä, T., Lönn, G., Schobesberger, S., Maso, M. D., Trimborn, A., Kulmala, M., and
- Worsnop, D.: Gas phase formation of extremely oxidized pinene reaction products in chamber and ambient air, Atmos Chem
   Phys, 12, 5113-5127, 2012.
- 471 Ehn, M., Thornton, J. A., Kleist, E., Sipila, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R., Lee, B.,
- 472 Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J.,
- 473 Nieminen, T., Kurten, T., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., Canagaratna, M., Dal Maso, M., Berndt, T., Petaja,
- 474 T., Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility
- 475 secondary organic aerosol, Nature, 506, 476-+, 10.1038/nature13032, 2014.
- 476 Eisele, F. L., and Tanner, D. J.: Measurement of the gas-phas concentration of H2SO4 and methane sulfonic-acid and
- estimates of H2SO4 production and loss in the atmosphere, J Geophys Res-Atmos, 98, 9001-9010, 10.1029/93jd00031,
  1993.
- 479 Frege, C., Ortega, I. K., Rissanen, M. P., Praplan, A. P., Steiner, G., Heinritzi, M., Ahonen, L., Amorim, A., Bernhammer,
- 480 A.-K., and Bianchi, F.: Influence of temperature on the molecular composition of ions and charged clusters during pure
- 481 biogenic nucleation, Atmos Chem Phys, 18, 65-79, 2018.
- 482 Hamed, A., Korhonen, H., Sihto, S. L., Joutsensaari, J., Järvinen, H., Petäjä, T., Arnold, F., Nieminen, T., Kulmala, M., and
- 483 Smith, J. N.: The role of relative humidity in continental new particle formation, Journal of Geophysical Research:
  484 Atmospheres, 116, 2011.
- Hanson, D., McMurry, P., Jiang, J., Tanner, D., and Huey, L.: Ambient pressure proton transfer mass spectrometry: detection
   of amines and ammonia, Environ Sci Technol, 45, 8881-8888, 2011.
- Hoffmann, T., O'Dowd, C. D., and Seinfeld, J. H.: Iodine oxide homogeneous nucleation: An explanation for coastal new
   particle production, Geophys Res Lett, 28, 1949-1952, 2001.
- 489 Hyttinen, N., Kupiainen-Määttä, O., Rissanen, M. P., Muuronen, M., Ehn, M., and Kurtén, T.: Modeling the charging of
- 490 highly oxidized cyclohexene ozonolysis products using nitrate-based chemical ionization, The Journal of Physical Chemistry
- 491 A, 119, 6339-6345, 2015.
- 492 Hyvönen, S., Junninen, H., Laakso, L., Maso, M. D., Grönholm, T., Bonn, B., Keronen, P., Aalto, P., Hiltunen, V., and Pohja,
- 493 T.: A look at aerosol formation using data mining techniques, Atmos Chem Phys, 5, 3345-3356, 2005.
- 494 Isaacman-VanWertz, G., Massoli, P., O'Brien, R., Lim, C., Franklin, J. P., Moss, J. A., Hunter, J. F., Nowak, J. B.,
- Canagaratna, M. R., and Misztal, P. K.: Chemical evolution of atmospheric organic carbon over multiple generations of
   oxidation, Nature chemistry, 1, 2018.
- Johnson, D., and Marston, G.: The gas-phase ozonolysis of unsaturated volatile organic compounds in the troposphere,
  Chemical Society Reviews, 37, 699-716, 2008.
- 499 Jokinen, T., Sipila, M., Junninen, H., Ehn, M., Lonn, G., Hakala, J., Petaja, T., Mauldin, R. L., Kulmala, M., and Worsnop, D.
- R.: Atmospheric sulphuric acid and neutral cluster measurements using CI-APi-TOF, Atmos Chem Phys, 12, 4117-4125,
  2012.
- 502 Jokinen, T., Sipilä, M., Richters, S., Kerminen, V. M., Paasonen, P., Stratmann, F., Worsnop, D., Kulmala, M., Ehn, M., and
- 503 Herrmann, H.: Rapid autoxidation forms highly oxidized RO2 radicals in the atmosphere, Angewandte Chemie International
- 504 Edition, 53, 14596-14600, 2014.





- Jonsson, Å. M., Hallquist, M., and Ljungström, E.: Impact of humidity on the ozone initiated oxidation of limonene,
   Δ3-carene, and α-pinene, Environ Sci Technol, 40, 188-194, 2006.
- 507 Jonsson, Å. M., Hallquist, M., and Ljungström, E.: Influence of OH scavenger on the water effect on secondary organic
- 508 aerosol formation from ozonolysis of limonene, Δ3-carene, and α-pinene, Environ Sci Technol, 42, 5938-5944, 2008.
- 509 Junninen, H., Ehn, M., Petaja, T., Luosujarvi, L., Kotiaho, T., Kostiainen, R., Rohner, U., Gonin, M., Fuhrer, K., Kulmala,
- 510 M., and Worsnop, D. R.: A high-resolution mass spectrometer to measure atmospheric ion composition, Atmos Meas Tech, 3,
- 511 1039-1053, 2010.
- 512 Kamens, R., Jang, M., Chien, C.-J., and Leach, K.: Aerosol Formation from the Reaction of α-Pinene and Ozone Using a
- 513 Gas-Phase Kinetics-Aerosol Partitioning Model, Environ Sci Technol, 33, 1430-1438, 10.1021/es980725r, 1999.
- 514 Kroll, J. H., Donahue, N. M., Jimenez, J. L., Kessler, S. H., Canagaratna, M. R., Wilson, K. R., Altieri, K. E., Mazzoleni, L.
- 515 R., Wozniak, A. S., Bluhm, H., Mysak, E. R., Smith, J. D., Kolb, C. E., and Worsnop, D. R.: Carbon oxidation state as a
- metric for describing the chemistry of atmospheric organic aerosol, Nature Chemistry, 3, 133-139, 10.1038/nchem.948,
  2011.
- 518 Kuang, C., McMurry, P. H., McCormick, A. V., and Eisele, F. L.: Dependence of nucleation rates on sulfuric acid vapor 519 concentration in diverse atmospheric locations, J Geophys Res, 113, 10.1029/2007jd009253, 2008.
- 520 Kuerten, A., Bergen, A., Heinritzi, M., Leiminger, M., Lorenz, V., Piel, F., Simon, M., Sitals, R., Wagner, A. C., and Curtius,
- 521 J.: Observation of new particle formation and measurement of sulfuric acid, ammonia, amines and highly oxidized organic
- 522 molecules at a rural site in central Germany, Atmos Chem Phys, 16, 12793-12813, 10.5194/acp-16-12793-2016, 2016.
- 523 Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V. M., Birmili, W., and McMurry, P. H.:
- 524 Formation and growth rates of ultrafine atmospheric particles: a review of observations, J Aerosol Sci, 35, 143-176,
- 525 10.1016/j.jaerosci.2003.10.003, 2004.
- 526 Kulmala, M., Petäjä, T., Ehn, M., Thornton, J., Sipilä, M., Worsnop, D., and Kerminen, V.-M.: Chemistry of atmospheric
- 527 nucleation: on the recent advances on precursor characterization and atmospheric cluster composition in connection with
- 528 atmospheric new particle formation, Annual review of physical chemistry, 65, 21-37, 2014.
- Kurten, A., Rondo, L., Ehrhart, S., and Curtius, J.: Calibration of a Chemical Ionization Mass Spectrometer for the
   Measurement of Gaseous Sulfuric Acid, J Phys Chem A, 116, 6375-6386, 2012.
- 531 Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Kurtén, T., Worsnop, D. R., and Thornton, J. A.: An iodide-adduct
- high-resolution time-of-flight chemical-ionization mass spectrometer: Application to atmospheric inorganic and organic
   compounds, Environ Sci Technol, 48, 6309-6317, 2014.
- 534 Li, Y., Poeschl, U., and Shiraiwa, M.: Molecular corridors and parameterizations of volatility in the chemical evolution of
- 535 organic aerosols, Atmos Chem Phys, 16, 3327-3344, 10.5194/acp-16-3327-2016, 2016.
- 536 Makkonen, R., Asmi, A., Kerminen, V. M., Boy, M., Arneth, A., Hari, P., and Kulmala, M.: Air pollution control and
- decreasing new particle formation lead to strong climate warming, Atmos Chem Phys, 12, 1515-1524,
  10.5194/acp-12-1515-2012, 2012.
- 539 McVay, R. C., Zhang, X., Aumont, B., Valorso, R., Camredon, M., La, Y. S., Wennberg, P. O., and Seinfeld, J. H.: SOA
- formation from the photooxidation of α-pinene: systematic exploration of the simulation of chamber data, Atmos Chem Phys, 16, 2785-2802, 2016.
- Merikanto, J., Spracklen, D., Mann, G., Pickering, S., and Carslaw, K.: Impact of nucleation on global CCN, Atmos Chem
   Phys, 9, 8601-8616, 2009.
- 544 Merikanto, J., Duplissy, J., Määttänen, A., Henschel, H., Donahue, N. M., Brus, D., Schobesberger, S., Kulmala, M., and
- 545 Vehkamäki, H.: Effect of ions on sulfuric acid water binary particle formation: 1. Theory for kinetic and nucleation -
- type particle formation and atmospheric implications, Journal of Geophysical Research: Atmospheres, 121, 1736-1751,2016.
- 548 Metzger, A., Verheggen, B., Dommen, J., Duplissy, J., Prevot, A. S., Weingartner, E., Riipinen, I., Kulmala, M., Spracklen, D.
- 549 V., Carslaw, K. S., and Baltensperger, U.: Evidence for the role of organics in aerosol particle formation under atmospheric





- 550 conditions, Proc Natl Acad Sci U S A, 107, 6646-6651, 10.1073/pnas.0911330107, 2010.
- 551 Molteni, U., Bianchi, F., Klein, F., Haddad, I. E., Frege, C., Rossi, M. J., Dommen, J., and Baltensperger, U.: Formation of
- highly oxygenated organic molecules from aromatic compounds, Atmos Chem Phys, 18, 1909-1921, 2018.
- 553 Napari, I., Noppel, M., Vehkamäki, H., and Kulmala, M.: Parametrization of ternary nucleation rates for H2SO4 NH3 -
- H2O vapors, Journal of Geophysical Research: Atmospheres, 107, 2002.
- 555 O'Dowd, C. D., Geever, M., Hill, M. K., Smith, M. H., and Jennings, S. G.: New particle formation: Nucleation rates and
- 556 spatial scales in the clean marine coastal environment, Geophys Res Lett, 25, 1661-1664, 1998.
- 557 Ortega, I. K., Donahue, N. M., Kurten, T., Kulmala, M., Focsa, C., and Vehkamaki, H.: Can Highly Oxidized Organics
- 558 Contribute to Atmospheric New Particle Formation?, J Phys Chem A, 120, 1452-1458, 10.1021/acs.jpca.5b07427, 2016.
- Pankow, J. F., and Asher, W. E.: SIMPOL. 1: a simple group contribution method for predicting vapor pressures and enthalpies of vaporization of multifunctional organic compounds, Atmos Chem Phys, 8, 2773-2796, 2008.
- 561 Poschl, U., Martin, S. T., Sinha, B., Chen, Q., Gunthe, S. S., Huffman, J. A., Borrmann, S., Farmer, D. K., Garland, R. M.,
- 562 Helas, G., Jimenez, J. L., King, S. M., Manzi, A., Mikhailov, E., Pauliquevis, T., Petters, M. D., Prenni, A. J., Roldin, P.,
- 563 Rose, D., Schneider, J., Su, H., Zorn, S. R., Artaxo, P., and Andreae, M. O.: Rainforest Aerosols as Biogenic Nuclei of
- 564 Clouds and Precipitation in the Amazon, Science, 329, 1513-1516, 2010.
- 565 Rissanen, M. P., Kurtén, T., Sipilä, M., Thornton, J. A., Kausiala, O., Garmash, O., Kjaergaard, H. G., Petäjä, T., Worsnop, D.
- R., and Ehn, M.: Effects of chemical complexity on the autoxidation mechanisms of endocyclic alkene ozonolysis products:
- 567 From methylcyclohexenes toward understanding α-pinene, The Journal of Physical Chemistry A, 119, 4633-4650, 2015.
- 568 Schobesberger, S., Junninen, H., Bianchi, F., Lonn, G., Ehn, M., Lehtipalo, K., Dommen, J., Ehrhart, S., Ortega, I. K.,
- 569 Franchin, A., Nieminen, T., Riccobono, F., Hutterli, M., Duplissy, J., Almeida, J., Amorim, A., Breitenlechner, M., Downard,
- 570 A. J., Dunne, E. M., Flagan, R. C., Kajos, M., Keskinen, H., Kirkby, J., Kupc, A., Kurten, A., Kurten, T., Laaksonen, A.,
- 571 Mathot, S., Onnela, A., Praplan, A. P., Rondo, L., Santos, F. D., Schallhart, S., Schnitzhofer, R., Sipila, M., Tome, A.,
- 572 Tsagkogeorgas, G., Vehkamaki, H., Wimmer, D., Baltensperger, U., Carslaw, K. S., Curtius, J., Hansel, A., Petaja, T.,
- 573 Kulmala, M., Donahue, N. M., and Worsnop, D. R.: Molecular understanding of atmospheric particle formation from sulfuric
- 574 acid and large oxidized organic molecules, Proc Natl Acad Sci U S A, 110, 17223-17228, 10.1073/pnas.1306973110, 2013.
- 575 Shen, X. J., Sun, J. Y., Zhang, Y. M., Wehner, B., Nowak, A., Tuch, T., Zhang, X. C., Wang, T. T., Zhou, H. G., Zhang, X. L.,
- 576 Dong, F., Birmili, W., and Wiedensohler, A.: First long-term study of particle number size distributions and new particle
- formation events of regional aerosol in the North China Plain, Atmos Chem Phys, 11, 1565-1580, 10.5194/acp-11-1565-2011,
  2011.
- 579 Sihto, S.-L., Kulmala, M., Kerminen, V.-M., Maso, M. D., Petäjä, T., Riipinen, I., Korhonen, H., Arnold, F., Janson, R., and
- 580 Boy, M.: Atmospheric sulphuric acid and aerosol formation: implications from atmospheric measurements for nucleation and
- early growth mechanisms, Atmos Chem Phys, 6, 4079-4091, 2006.
- 582 Thomson, B., and Iribarne, J.: Field induced ion evaporation from liquid surfaces at atmospheric pressure, The Journal of
- 583 Chemical Physics, 71, 4451-4463, 1979.
- 584 Tobias, H. J., Docherty, K. S., Beving, D. E., and Ziemann, P. J.: Effect of relative humidity on the chemical composition of
- 585 secondary organic aerosol formed from reactions of 1-tetradecene and O3, Environ Sci Technol, 34, 2116-2125, 2000.
- 586 Tobias, H. J., and Ziemann, P. J.: Kinetics of the gas-phase reactions of alcohols, aldehydes, carboxylic acids, and water with
- the C13 stabilized Criegee intermediate formed from ozonolysis of 1-tetradecene, The Journal of Physical Chemistry A, 105,
   6129-6135, 2001.
- 589 Tröstl, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Ahlm, L., Frege, C., Bianchi, F., and Wagner, R.:
- 590 The role of low-volatility organic compounds in initial particle growth in the atmosphere, Nature, 533, 527-531, 2016.
- 591 Vehkamäki, H., and Riipinen, I.: Thermodynamics and kinetics of atmospheric aerosol particle formation and growth,
- 592 Chemical Society Reviews, 41, 5160-5173, 2012.
- 593 Weber, R., Marti, J., McMurry, P., Eisele, F., Tanner, D., and Jefferson, A.: Measurements of new particle formation and
- <sup>594</sup> ultrafine particle growth rates at a clean continental site, Journal of Geophysical Research: Atmospheres, 102, 4375-4385,

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595 1997.

- 596 Weber, R., McMurry, P. H., Mauldin, R., Tanner, D., Eisele, F., Clarke, A., and Kapustin, V.: New particle formation in the
- remote troposphere: A comparison of observations at various sites, Geophys Res Lett, 26, 307-310, 1999.
- 598 Zhang, R., Khalizov, A., Wang, L., Hu, M., and Xu, W.: Nucleation and growth of nanoparticles in the atmosphere, Chem
- 599 Rev, 112, 1957-2011, 10.1021/cr2001756, 2012.
- 600 Zhang, X., McVay, R. C., Huang, D. D., Dalleska, N. F., Aumont, B., Flagan, R. C., and Seinfeld, J. H.: Formation and
- evolution of molecular products in α-pinene secondary organic aerosol, Proceedings of the National Academy of Sciences,
  112, 14168-14173, 2015.
- 603 Zhao, J., Khalizov, A., Zhang, R., and McGraw, R.: Hydrogen-bonding interaction in molecular complexes and clusters of
- aerosol nucleation precursors, The Journal of Physical Chemistry A, 113, 680-689, 2009.
- 605 Zhao, J., Eisele, F. L., Titcombe, M., Kuang, C., and McMurry, P. H.: Chemical ionization mass spectrometric measurements
- of atmospheric neutral clusters using the cluster CIMS, Journal of Geophysical Research: Atmospheres, 115, 2010.
- 607 Zhao, J., Ortega, J., Chen, M., and Mcmurry, P. H.: Dependence of particle nucleation and growth on high molecular weight
- 608 gas phase products during ozonolysis of  $\alpha$ -pinene, Atmospheric Chemistry & Physics, 13, 7631-7644, 2013.





## 610 Table 1. Experiment conditions and products.

Precursor	Exp (#)	Monoterpene (ppbv)	O <sub>3</sub> (ppbv)	Cyclo-hexane (ppmv)	Initial rate <sup>1</sup> (10 <sup>8</sup> molecules cm <sup>-3</sup> s <sup>-1</sup> )	O <sub>3</sub> consumption <sup>2</sup> (ppb)	SOA <sup>3</sup> (μg m <sup>-3</sup> )
Limonene	1	1085	900±10	0	1410	159-166	138-208
	2	1085	900±10	217	1410	139-150	81-147
	3	54	350±5	0	27.3	34-41	0
α-pinene	4	1111	900±10	0	625	103-110	761-1042
	5	1111	900±10	222	625	93-102	414-735
	6	54	350±5	0	11.8	23-30	0
$\triangle^3$ -carene	7	1111	900±10	0	267	72-89	55-93
	8	1111	900±10	222	267	70-86	34-92
	9	54	350±5	0	5.05	11-16	0

611 <sup>1</sup>At room temperature (298K), the rate coefficients for limonene,  $\alpha$ -pinene and  $\Delta^3$ -carene to react with  $O_3$  were 200×10<sup>-18</sup>,

612 **86.6×10<sup>-18</sup>, 37×10<sup>-18</sup> cm<sup>3</sup> molecule**<sup>-1</sup> s<sup>-1</sup>, respectively.

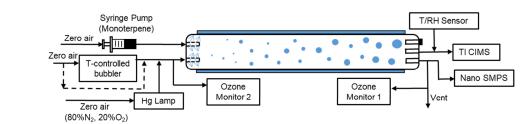
613 <sup>2</sup>O<sub>3</sub> consumption values were calculated from the difference between inlet and outlet O<sub>3</sub> concentrations.

614 <sup>3</sup>SOA mass concentrations were calculated from SMPS-measured volume concentrations and an assumed organic effective density

615 (**1.2 g cm**<sup>-3</sup>).







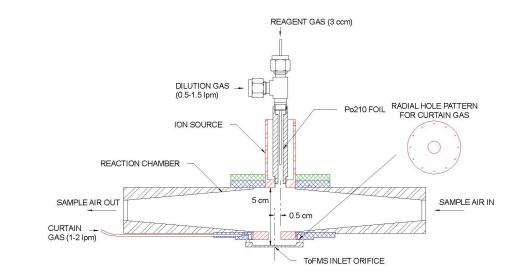
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618 Figure 1. Experiment setup for the flow tube experiments. The 8.5 L flow tube was placed at a temperature-controlled room

619 (21±1°C) and covered. The total flowrate was 8.5 LPM. The RH was adjusted by mixing temperature controlled bubbler flow with

620 dry zero air.

621



623 Figure 2. Schematic of the transverse ionization (TI) inlet, showing the N<sub>2</sub> curtain gas configuration. The relative position of the

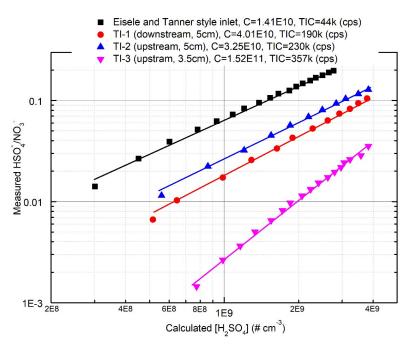
624 ion source to the inlet orifice is adjustable. The configuration shown here is the most sensitive in calibrations with H<sub>2</sub>SO<sub>4</sub> (see

625 Section 3.1).

626







627

Figure 3. Comparison of the sensitivities for the two inlets to  $H_2SO_4$ . The calibration process followed that reported by Kurten et al. (2012) and is discussed in detail in the supplementary material. TI-1, 2, 3 represent different locations of the ion source relative to the inlet orifice of the mass spectrometer. "upstream" and "downstream" indicated 0.5 cm upstream or downstream along the sample flow axis and "3.5 cm" and "5 cm" indicate 3.5 or 5 cm away from the inlet orifice along the reagent ion flow axis.

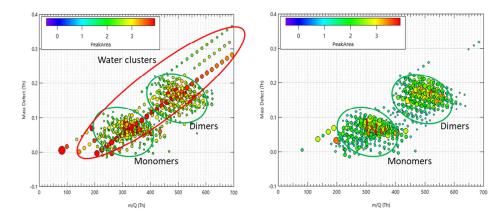




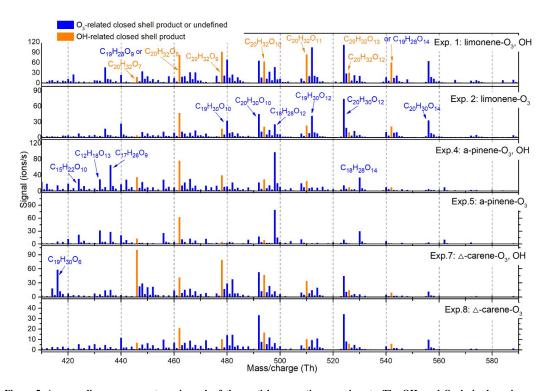
Figure 4. Mass defect plots of α-pinene ozonolysis HOMs with 0 LPM (left) and 1 LPM (right) N<sub>2</sub> curtain gas flow when RH> 85%,
with monomer and dimer HOMs circled in green. The most intense ions comprising 60% of the total ion count are plotted for

635 clarity. H<sub>2</sub>O clusters (H<sub>2</sub>O)<sub>m</sub>(HNO<sub>3</sub>)<sub>n</sub>NO<sub>3</sub><sup>-</sup> (m=1~30, n=0~2) are circled in red in the left plot and are notably absent with the

 $\label{eq:solution} 636 \qquad \text{application of the } N_2 \, \text{curtain gas.} \, (\text{HNO}_3)_2 \text{NO}_3 \, \text{and } \text{HNO}_3 \text{NO}_3 \, \text{are much more likely to cluster with } H_2 \text{O than } \text{NO}_3 \, \text{.}$ 



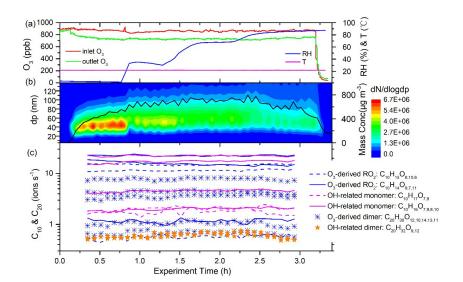




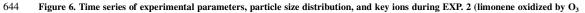
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Figure 5. Average dimer mass spectrum in each of the particle generation experiments. The OH- and O<sub>3</sub>-derived species were
distinguished by comparing relative abundance of experiments with and without OH scavenger. All the peaks shown were in the
form of adducts with NO<sub>3</sub> or HNO<sub>3</sub>NO<sub>3</sub> reagent ions.

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643



645 without OH scavenger). (a) inlet and outlet O<sub>3</sub> concentrations, temperature, and RH; (b) Particle size distribution and integrated



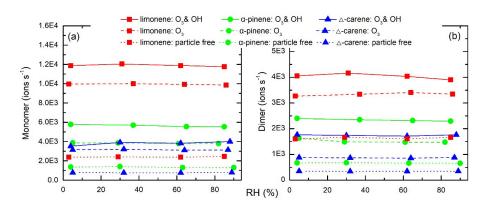


646 mass concentrations (assuming effective density is 1.2 g cm<sup>-3</sup>); (c) Some of the main HOMs detected by TI-CIMS with NO<sub>3</sub><sup>-</sup> reagent

647 ion. The subscript oxygen numbers in the formulae were ranked (left-to-right) according to signal abundance of the corresponding

648 molecule.

649

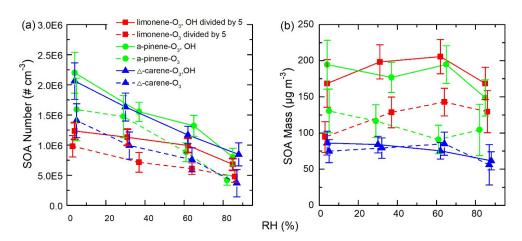




651 Figure 7. Average (a) monomer and (b) dimer HOMs signal intensity (ions s<sup>-1</sup>) as a function of RH in each experiment. Monomer

652 signals were the sum of  $C_{5^{-1}0}$  molecules and dimer signals were the sum of  $C_{15^{-2}0}$  molecules. No obvious signal change was seen for

- 653 increasing RH in any of the experiments.
- 654



655

656 Figure 8. SOA (a) number and (b) mass concentrations as a function of RH during different experiments. The error bars were

657 calculated using both the statistical errors of all individual size distributions during each RH stage and assuming a systematic CPC

<sup>658</sup> counting error of 10%).





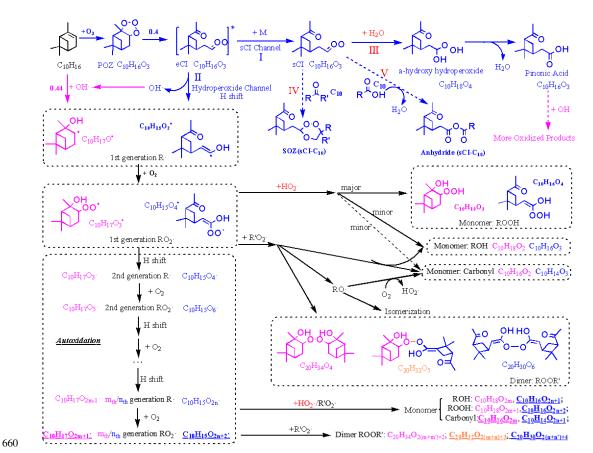


Figure 9. Proposed key steps in the formation of the representative  $C_{10}$  and  $C_{20}$  closed shell products from *a*-pinene oxidation and possible water vapor influence. Dashed lines represent pathways that may or may not happen, depending on the situation. Pink and blue colors represented the pathways or products from O<sub>3</sub> and OH oxidation, respectively. Common pathways or products are indicated in black type. Orange colors represented the combined products of O<sub>3</sub> and OH chemistry. Red colors highlight the direct or indirect influence of water. Underlined formulae were the main products observed from the mass spectrum.  $C_{10}H_{18}O_{2m}$  and  $C_{20}H_{34}O_{2(m+m')}$  were not observed in our spectrum, but they dominated the spectrum in other reported experiments where extra OH was generated (Berndt et al., 2016).





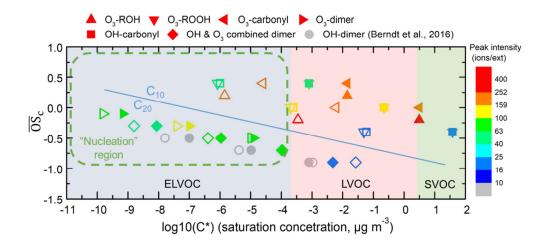




Figure 10. Vapor saturation mass concentration C\* (T=298 K) of the major C<sub>10</sub> and C<sub>20</sub> closed shell products were predicted with
SIMPOL.1 (open points) (Pankow and Asher, 2008) and Molecular Corridor method (filled points) (Li et al., 2016). O<sub>3</sub>-derived,
OH-related and OH-derived monomers and dimers are presented in different shapes. The peak intensity, represented by color, is

673 from Exp.1 (limonene oxidation without OH scavenger). The gray points, which represent OH-derived dimers, are dominating

674 products in OH initiated oxidation experiments (Berndt et al., 2016) while not observed in this study. Data used in this figure are

675 given in Table S1. The nucleation region is from Donahue et al. (2013).