1	Retrospective analysis of 2015-2017 winter-time PM _{2.5} in China: response to emission
2	regulations and the role of meteorology
3 4	
5	Dan Chen ^{1*} , Zhiquan Liu ^{2*} , Junmei Ban ² , Pusheng Zhao ¹ , and Min Chen ¹
6 7	¹ Institute of Urban Meteorology, China Meteorological Administration, Beijing, 100089, China ² National Center for Atmospheric Research, Boulder, CO, 80301, USA
8	
9	
10	February, 2019
11	

^{*} Corresponding author: Dr. Zhiquan Liu (<u>liuz@ucar.edu</u>) and Dr. Dan Chen (dchen@ium.cn)

1 Abstract

To better characterize anthropogenic emission-relevant aerosol species, the GSI-WRF/Chem data 2 assimilation system was updated from the GOCART aerosol scheme to the MOSAIC-4BIN scheme. Three 3 years (2015-2017) of wintertime (January) surface PM_{2.5} observations from 1600+ sites were assimilated 4 hourly using the updated 3DVAR system. In the control experiment (without assimilation) using 2010 MEIC 5 emissions, the modeled January averaged PM_{2.5} concentrations were severely overestimated in the Sichuan 6 Basin, Central China, Yangtze River Delta, and Pearl River Delta by 98-134, 46-101, 32-59, and 19-60 µg m⁻ 7 ³, respectively, indicating that the emissions for 2010 are not appropriate for 2015-2017, as strict emission 8 control strategies were implemented in recent years. Meanwhile, underestimations of 11-12, 53-96, and 22-40 9 ug m⁻³ were observed in northeastern China, Xinjiang and the Energy Golden Triangle, respectively. The 10 assimilation experiment significantly reduced both high and low biases to within $\pm 5 \ \mu g \ m^{-3}$. 11

The observations and the reanalysis data from the assimilation experiment were used to investigate the 12 year-to-year changes and the driving factors. The role of emissions was obtained by subtracting the 13 meteorological impacts (by control experiments) from the total combined differences (by assimilation 14 experiments). The results show a reduction in $PM_{2.5}$ of approximately 15 µg m⁻³ for the month of January from 15 2015 to 2016 in the North China Plain (NCP), but meteorology played the dominant role (contributing a 16 reduction of approximately 12 µg m⁻³). The change (for January) from 2016 to 2017 in NCP was different; 17 meteorology caused an increase in $PM_{2.5}$ of approximately 23 µg m⁻³, while emission control measures caused 18 a decrease of 8 μ g m⁻³, and the combined effects still showed a PM_{2.5} increase for that region. The analysis 19 confirmed that emission control strategies were indeed implemented and emissions were reduced in both years. 20 Using a data assimilation approach, this study helps identify the reasons why emission control strategies may 21 or may not have an immediately visible impact. There are still large uncertainties in this approach, especially 22 the inaccurate emission inputs, and neglecting aerosol-meteorology feedbacks in the model can generate large 23 uncertainties in the analysis as well. 24

25 1. Introduction

Anthropogenic $PM_{2.5}$ (fine particulate matter with an aerodynamic diameter smaller than 2.5 μ m) is known as a robust indicator of mortality and other negative health effects associated with ambient air pollution.

PM_{2.5} components are originate not only from primary emissions but also from secondary formations through 1 various precursors (e.g., SO₂, NO_x, and VOCs). Regional haze with extremely high PM_{2.5} concentrations 2 (exceeding the WHO standard tenfold) has become the primary air quality concern in China, especially over 3 northern China (e.g., Wang et al. 2014a, 2014b; Han et al. 2015; Sun et al. 2015). To control PM_{2.5} pollution 4 and improve the overall air quality, a series of strict pollution control strategies have been implemented by the 5 government since 2010, including the Guiding Options on Promoting the Joint Prevention and Control of Air 6 Pollution to Improve Regional Air Quality (The Central Government of the People's Republic of China, 2010) 7 and the Atmospheric Pollution Prevention and Control Action Plan (The Central Government of the People's 8 Republic of China, 2013), in which the government stated that environmental-related equipment (for flue-gas 9 desulfurization, selective catalyst reduction, exhaust dust removal, etc.) are mandatory for both industries and 10 vehicles. In addition to long-term pollution control strategies, different emergency measures under different 11 pollution alerts were also implemented occasionally. For example, the production of large industrial sources 12 (coal burning and cement) was limited to reduce emissions, construction sites were restricted to prevent 13 fugitive dust pollution, and traffic restrictions were implemented on even- and odd-numbered license plates. 14 These emission control strategies were stricter and implemented more often in northern China than anywhere 15 else in winter, when haze events occur more frequently. These control strategies were expected to reduce both 16 the concentrations of significant precursors (e.g., SO_2 , NO_x) and the emissions of $PM_{2.5}$. 17

Despite these strict emission control strategies, the ambient PM_{2.5} concentrations in major cities still 18 fluctuated during the wintertime from year to year. For example, the overall January PM_{2.5} concentrations in 19 74 cities generally decreased from 2015 to 2016, but the concentrations in January 2017 were still higher than 20 (Ambient *Monthly* 2015-01/2016-01/2017-01, those in 2016 Air Quality Report 21 http://www.cnemc.cn/kgzlzkbgyb2092938.jhtml). While annual emission reduction trends were expected 22 from 2015 to 2017, the overall increase in the surface concentrations observed in January 2017 contradicted 23 these expectations, thereby indicating that other factors (especially meteorological conditions) in addition to 24 emissions may play important roles. Some studies have attempted to investigate the variability of air pollution 25

and the effects of climate change on wintertime air pollution by using statistical data. Li *et al.* (2016) indicated that wintertime fog-haze days across central and eastern China have a close relationship with the East Asian winter monsoon. Zuo *et al.* (2015) concluded that the significant weakening and strengthening of the Siberian high and East Asian trough, respectively, are the two main factors for the occurrence of extreme cold and extreme warm events over China in winter, when warm air boosts air pollution. In addition to utilizing statistical methodology, it is necessary to distinguish the roles of emissions and meteorology to further investigate the driving factors of interannual air pollution changes.

Regional air quality models are important tools for scientifically understanding the formation of haze 8 events, technically constructing forecasts, and evaluating the effects of control strategies. For regional 9 modeling studies, emission inventories are important for reflecting the emission inputs into the atmosphere. 10 Generally, an emission inventory is based on a "bottom-up" methodology, thereby relying on the statistics of 11 energy activity and emission factors, etc. However, uncertainties in energy statistics can cause variations in 12 the emission estimates (Zhao et al., 2017; Hong et al., 2017; Zhi et al., 2017). For regional modeling 13 applications, the total emissions based on statistics are spatially and temporally distributed according to 14 relevant factors (He, 2012). Nevertheless, the occasional emission control strategies implemented in winter 15 can cause large uncertainties in not only total emission estimations but also spatial and temporal allocations, 16 which would lead to large biases in the model simulations. 17

In addition to the uncertainties in emission inventories, deficiencies in the model chemistry can also cause 18 model uncertainties. Increasing numbers of observations have revealed that anthropogenic emission-relevant 19 aerosol species, such as sulfate, nitrate and ammonium (denoted as SNA), are the predominant inorganic 20 species in the wintertime PM_{2.5} in China (Wang et al., 2014c; Yang et al., 2015). Various reaction paths during 21 haze events have also been proposed (e.g. Zheng et al., 2015; Cheng et al., 2016; Wang et al., 2016; Li et al., 22 2017; Moch et al., 2018; Wang et al., 2018; Shao et al., 2019). For example, Moch et al. (2018) used a 1-D 23 model and revealed the importance of aqueous-phase chemistry of HCHO and S(IV) in cloud droplets by 24 forming a S(IV)-HCHO adduct, hydroxymethane sulfonate. Shao et al. (2019) implemented four 25

heterogeneous sulfate formation mechanisms (via H₂O₂, O₃, NO₂, and transition metal ions on aerosols) into 1 GEOS-Chem model which partially reduced the modeled low bias in sulfate concentrations. However, a 2 scientific consensus regarding the importance of the reaction paths has not yet been reached partially due to 3 the uncertainties of aerosol liquid water content, pH, and ionic strength etc. The original WRF/Chem model 4 with either the Goddard Chemistry Aerosol Radiation and Transport (GOGART, Chin et al., 2000, 2002) or 5 the Model for Simulating Aerosol Interactions and Chemistry (MOSAIC)-4BIN aerosol scheme failed to 6 reproduce the highest PM_{2.5} concentrations; it is assumed that this failure is due to missing 7 heterogeneous/aqueous reactions. In Chen et al. (2016, hereafter Chen16), we added three heterogeneous 8 reactions (SO₂-to-H₂SO₄ and NO₂/NO₃-to-HNO₃) to the WRF/Chem model based on the MOSAIC-4BIN 9 aerosol scheme. Although the reaction paths may still not be comprehensively understood, the new MOSAIC-10 4BIN aerosol scheme significantly improved the simulation of sulfate, nitrate, and ammonium on polluted 11 days in terms of the concentrations of those species and their partitioning. 12

Data assimilation (DA), that is, the combination of observations with numerical model output, has proven 13 to be skillful at improving aerosol forecasts (e.g., Collins et al., 2001; Pagowski et al., 2010; Liu et al., 2011; 14 Liu et al., 2016; Zhang et al., 2016). Liu et al. (2011, hereafter Liu11) implemented DA on AOD estimates 15 within the National Centers for Environmental Prediction (NCEP) gridpoint statistical interpolation (GSI) 16 three-dimensional variational (3DVAR) DA system coupled with the GOCART aerosol scheme within the 17 Weather Research and Forecasting/Chemistry (WRF/Chem) model (Grell et al., 2005). Schwartz et al. (2012, 18 hereafter S12) and Jiang et al. (2013, hereafter Jiang13) extended the above system to assimilate surface PM_{2.5} 19 and PM₁₀. The evaluation results demonstrated improved aerosol forecasts from the DA system in studies over 20 East Asia and the United States. 21

Following Liu11, S12 and Chen16, we updated the GSI-WRF/Chem system by changing from the GOCART aerosol scheme to the MOSAIC-4BIN aerosol scheme to better characterize the complex $PM_{2.5}$ pollution in China. We applied the updated system to assimilate $PM_{2.5}$ concentrations of January 2015, 2016 and 2017 for two purposes: 1) to reproduce the $PM_{2.5}$ output by the DA system and 2) to investigate the

different impacts of meteorological conditions and emissions on the PM_{2.5} pollution in different years. In this 1 paper, section 2 provides descriptions of the model, observations and methodology and addresses the updated 2 GSI-WRF/Chem-coupled DA system with the MOSAIC-4BIN aerosol scheme. In section 3, the assimilation 3 results for the PM_{2.5} concentrations from January 2015, 2016 and 2017 are presented and compared with 4 surface observations (PM_{2.5} total mass) to evaluate the DA system. In contrast to previous applications 5 emphasizing the forecast skill improvement achieved by the DA system, we fully utilized reanalysis data to 6 investigate the driving factors of pollution and to differentiate the roles played by meteorological conditions 7 and emissions in different years by analyzing the reanalysis data and model simulations. The results are given 8 in section 4, and the conclusions are given in section 5. 9

10 2. Model description, observations and methodology

The WRF/Chem settings are very similar to those of Chen16, although Chen16 focused on the SNA 11 aerosols in the North China Plain during October 2014; in addition, several heterogeneous reactions were 12 newly added to the original chemistry modules to improve the SNA simulation performance. The DA system 13 used herein was based upon the NCEP GSI system extended by Liu11 and S12. We assimilated surface PM_{2.5} 14 observations, and the only difference is that the MOSAIC-4Bin aerosol scheme (32 PM species) was chosen 15 for the WRF/Chem model instead of the GOCART aerosol scheme. Thus, the 3-D mass mixing ratios of those 16 MOSAIC species at each grid point composed the analysis (or control) variables in the GSI 3DVAR 17 minimization process. 18

Here, only a brief summary of the WRF/Chem configuration is provided below prior to a description of
the updated GSI DA system and the settings used in this work. The most important differences are noted, e.g.,
the forward operator for observations in the GSI system.

22 2.1 WRF/Chem model and emissions

As in Chen16, version 3.6.1 of the WRF/Chem model was used in this study (Grell *et al.*, 2005; Fast *et al.*, 2006). The physical parameterizations employed in the WRF/Chem model were identical to those of

1	Chen16, and they are listed in Table 1. The Carbon-Bond Mechanism version Z (CBMZ) and the Model for
2	Simulating Aerosol Interactions and Chemistry (MOSAIC) were used as the gas phase and aerosol chemical
3	mechanisms, respectively, in this study. The aerosol species in MOSAIC are defined as black carbon (BC),
4	organic compounds (OC), sulfate (SO ₄), nitrate (NO ₃), ammonium (NH ₄), sodium (NA), chloride (CL) and
5	other inorganic compounds (OIN). We used 4 size bins with aerosol diameters ranging from 0.039-0.1, 0.1-
6	1.0, 1.0-2.5, and 2.5-10 μ m. The 24 variables in the first three bins (8 species times 3 bins) consist of the PM _{2.5}
7	total. The newly added relative humidity (RH)-dependent SO ₂ -to-H ₂ SO ₄ and NO ₂ /NO ₃ -to-HNO ₃
8	heterogeneous reactions (details are provided in Chen16) were also applied in the simulations.
9	The model domain with a 40-km horizontal grid spacing covers most of China and the surrounding
10	regions (Fig. 2), and there are 57 vertical levels extending from the surface to 10 hPa. The simulation started
11	from Dec. 20 of the previous year; the first eleven days were treated as a spin-up period and were not used in

- 12 our analyses.
- 13

Table 1. WRF/Chem model configuration

Aerosol scheme	MOSAIC (4 bins) (Zaveri et al., 2008)
Photolysis scheme	Fast-J (Wild et al., 2000)
Gas phase chemistry	CBM-Z (Zavier et al., 1999)
Cumulus parameterization	Grell 3D scheme
Short-wave radiation	Goddard Space Flight Center Shortwave radiation scheme (Chou and Suarez, 1994)
Long-wave radiation	RRTM (Mlawer et al., 1997)
Microphysics	Single-Moment 6-class scheme (Grell and Devenyi, 2002)
Land-surface model	NOAH LSM (Chen and Dudhia, 2001)
Boundary layer scheme	YSU (Hong <i>et al.</i> , 2006)
Meteorology initial and boundary conditions	GFS analysis and forecast every 6 hour
Initial condition for chemical species	11-day spin-up
Boundary conditions for chemical species	averages of mid-latitude aircraft profiles (McKeen et al., 2002)
Dust and sea salt Emissions	GOCART

As in Chen16, the Multi-resolution Emission Inventory for China (MEIC) (Zhang *et al.*, 2009; Lei *et al.*,

15 2011; He 2012; Li et al., 2014) for January 2010 was used as the emission input, as it is the only emission

1	inventory that was publicly available when the study was conducted. The original grid spacing of the MEIC
2	is $0.25^{\circ} \times 0.25^{\circ}$, and this inventory has been processed to match the model grid spacing (40 km). The spatial
3	distributions of primary PM _{2.5} , SO ₂ , NO _x and NH ₃ emissions are shown in Fig. 2. The MEIC-2010 emission
4	inventory has already been applied in other studies (e.g., Wang et al., 2014a; Zheng et al., 2015) for
5	simulations over China in the past few years; these recent studies found that the MEIC provides reasonable
6	estimates of total emissions but is subject to uncertainties in the spatial allocations of these emissions over
7	small spatial scales. For our simulations, uncertainties may also arise from two other sources: the difference
8	between the emission base year (2010) and our simulation period (2015 through 2017) and the monthly
9	allocations. As the Chinese government has implemented strict control strategies to ensure an improved air
10	quality during the winter season since 2013, significant reductions in emissions, including primary PM and
11	precursor compounds (SO ₂ and NO _x), in regions with the strict implementation of these policies relative to the
12	year 2010 are expected for our simulation period. A reduction in SO ₂ pollution of approximately 50% was
13	observed from 2012-2015 for the North China Plain from OMI satellite data (Krotkov et al., 2016). National
14	anthropogenic emission reductions of approximately 67%, 17%, and 35% from 2012-2017 for SO ₂ , NO _x , and
15	PM _{2.5} , respectively, were assumed by the bottom-up EI methodology (Zheng et al., 2018). However, the
16	expansion and relocation of the energy industry caused emission increases in northwestern China (Ling et al.,
17	2017). In addition, the uncertainties of allocated emissions in the winter season will be much larger than those
18	in other seasons. For example, Zhi et al. (2017) conducted a village energy survey and revealed an enormous
19	discrepancy in the amount of rural raw coal used for winter heating in northern China, implying an extreme
20	underestimation of rural household coal consumption by the China Energy Statistical Yearbooks. These
21	changes and uncertainties of emissions in the model would introduce errors into the NO_DA simulation.
22	However, the inhomogeneous spatial changes and large uncertainties of seasonal allocations made it difficult
23	to simply scale the original emission inventory for our study period.

24 2.2 Updated GSI 3DVAR DA system

25 The NCEP's GSI 3DVAR DA system was used to assimilate surface PM_{2.5} observations. The GSI 3DVAR

DA system calculates a best-fit analysis considering the observations (hourly surface PM_{2.5} concentrations in our case) and background fields (a 1-hr short-term WRF/Chem forecast in our case) weighted by their error characteristics. The GSI 3DVAR DA system produces an analysis in a model grid space through the minimization of a scalar objective function J(x) given by

5
$$J(\mathbf{x}) = \frac{1}{2} (\mathbf{x} - \mathbf{x}_b)^{\mathrm{T}} \mathbf{B}^{-1} (\mathbf{x} - \mathbf{x}_b) + \frac{1}{2} [H(\mathbf{x}) - \mathbf{y}]^{\mathrm{T}} \mathbf{R}^{-1} [H(\mathbf{x}) - \mathbf{y}],$$
(1)

6 where $\mathbf{x}_{\mathbf{b}}$ denotes the background vector (with dimension m), \mathbf{y} is a vector of observations (with dimension 7 p), and **B** and **R** represent the background and observation error covariance matrices of dimensions m × m 8 and p × p, respectively. The covariance matrices determine the relative contributions of the background and 9 observation terms to the final analysis. *H* is the potentially nonlinear "observation operator" that interpolates 10 the model grid point values into observation spaces and converts model-predicted variables into observed 11 quantities.

12 2.2.1 PM_{2.5} observation operator

In our updated DA system, GSI was used to assimilate surface $PM_{2.5}$ total mass observations, whereas the WRF/Chem model predicts the $PM_{2.5}$ total mass as different prognostic variables depending on the chosen aerosol scheme. As we chose the MOSAIC-4Bin aerosol scheme, the analysis variables here were the 3D mass mixing ratios of the 24 MOSAIC aerosol variables at each grid point. The model-simulated $PM_{2.5}$ observations $M_{PM_{2.5}}$ were computed by summing the 24 species as

18
$$M_{PM_{25}} = \sum_{i=1}^{3} [BC_i + OC_i + SO4_i + NO3_i + NH4_i + CL_i + NA_i + OIN_i],$$
(2)

where *i* denotes the bin number in the MOSAIC aerosol scheme, where the first three bins consist of the PM_{2.5} total, and BC, OC, SO₄, NO₃, NH₄, NA, CL, and OIN denote black carbon, organic compounds, sulfate, nitrate, ammonium, sodium, chloride and other inorganic compounds, respectively. This formula is identical to that used in the WRF/Chem MOSAIC scheme to diagnose PM_{2.5}. The WRF-Chem-simulated aerosol mixing ratios of the species listed inside the brackets of Eq. 2 are in units of $\mu g k g^{-1}$, and thus, the dry air density ρ_d is multiplied to convert the units into $\mu g m^{-3}$ for consistency with the observations.

Since only surface PM_{2.5} total mass observations were assimilated to analyze the 3D mass mixing ratios 1 of 24 aerosol variables, the 3DVAR problem was underconstrained. Due to the lack of species and vertical 2 information provided by the observations, the only mathematical solution is to utilize prior information from 3 the model background. In the GSI system, the distribution of the analysis increments (the difference between 4 the analysis and background) onto the different species was mostly model driven with the observation and 5 background error covariance matrices acting as the main constraints. This speciated approach to aerosol DA 6 within a variational system was introduced by Liu11 and further applied by S12 and Jiang13. By using 7 individual aerosol species as the control variables, no assumptions were made regarding the contribution of 8 each species' mass to the total aerosol mass or to the shapes of the vertical profiles. 9

10 2.2.2 PM_{2.5} observations and errors

Hourly surface PM_{2.5} observations for January 2015-2017 were obtained from the China National Environmental Monitoring Center (CNEMC). There are 1600+ sites in our modeling domain. As the 1600+ monitoring sites fall into 531 model grids, all observations within the same grid are averaged (as well as the latitude and longitude) for the purpose of performing statistical calculations and evaluation. The observation sites (Fig. 3) span mostly northern, central and eastern China, while the sites are relatively sparse in western China.

The observation error covariance matrix \mathbf{R} in Eq. (1) contains both measurement and representativeness 17 errors. Pagowski *et al.* (2010) used a measurement error (ε_0) of 2 µg m⁻³. To associate higher PM_{2.5} values 18 with larger measurement errors, S12 defined the measurement error as $\varepsilon_0 = 1.5 + 0.0075 \times M_{PM_2 \text{ s}}$, where 19 $M_{PM_{2.5}}$ denotes an AIRNow PM_{2.5} observation and the units of each term are $\mu g m^{-3}$. According to the PM_{2.5} 20 Auto-Monitoring Instrument Technical Standard and Requirement (China National Environmental Monitoring 21 Center, 2013), three continuous online monitoring methods, namely, a beta-ray plus dynamic heating system, 22 a beta-ray plus dynamic heating system plus light scattering system, and a tapered element oscillating 23 microbalance plus filter dynamic measurement system, are used at the national monitoring sites to satisfy the 24

1 requirements that the display resolution should be less than 1 μ g m⁻³ and the error should be less than 2 5μ g m⁻³(within 24 hours). To reflect the confidence in the hourly observations, the measurement error ε_0 in 3 this study is defined as $\varepsilon_0 = 1.0 + 0.0075 \times M_{PM_{2.5}}$, where $M_{PM_{2.5}}$ denotes a PM_{2.5} observational value 4 (unit: μ g m⁻³).

Representativeness errors reflect the inaccuracies in the forward operator and in the interpolation from
the model grid to the observation location. Elbern *et al.* (2007), Pagowski *et al.* (2010), S12 and Jiang13
defined the representativeness error (ε_r) as

$$\varepsilon_r = \gamma \varepsilon_0 \sqrt{\frac{\Delta x}{L}},\tag{3}$$

9 where γ is an adjustable parameter scaling ε_0 ($\gamma = 0.5$ was used here), Δx is the grid spacing (40 km in 10 our case) and L is the radius of influence of an observation (set to 2 km for urban sites). These parameter

settings were based on the performance of sensitivity tests. The total PM_{2.5} error ($\epsilon_{PM2.5}$) is defined as

$$\varepsilon_{PM2.5} = \sqrt{\varepsilon_0^2 + \varepsilon_r^2},\tag{4}$$

which constituted the diagonal elements in the **R** matrix. The PM_{2.5} data were provided in near-real time without any data quality control. To ensure the data quality before DA, PM_{2.5} observational values larger than 15 1000 μ g m⁻³ (the maximum display limit of the monitoring system) were deemed unrealistic in the filter 16 process and thus were not assimilated. In addition, observations leading to innovations/deviations 17 (observations minus the model-simulated values determined from the first-guess fields) exceeding 18 500 μ g m⁻³ were also omitted for the stability of the DA optimization step.

19 **2.2.3 Background error covariance**

8

12

Similar to Jiang13, the background error covariance (BEC) statistics for each analysis variable required by the 3DVAR algorithm were computed by utilizing the NMC method (Parrish and Derber, 1992) based upon the one-month WRF/Chem forecast for January 2015. No cross-correlation between different species was considered. The standard deviations and horizontal/vertical correlation length scales of the background errors (separated for each aerosol species) were calculated using the method described by Wu *et al.* (2002). These

data were used as constraints for the distributions of the PM components. It is important to have phenomena-1 specific background error statistics to allow for an appropriate adjustment of individual species. The domain-2 averaged standard deviations of the background errors for 6 aerosol species (BC, OC, SO₄, NO₃, NH₄, and 3 OIN) in the first three size bins are shown in Fig. 1 as a function of the vertical model level; CL and NA are 4 not shown here because they are excessively small relative to the other PM species. By using the MOSAIC 5 aerosol scheme, the characteristics of different aerosol species in different size bins are more appropriate for 6 the China region in the model. As shown in Fig. 1, the standard deviations of different aerosol species errors 7 are different in the three size bins; the errors of NO₃, OIN and SO₄ are relatively larger than those of the other 8 species in the three size bins; OC is also important, especially in the second (0.1-1.0 µm) and third (1.0-9 2.5 µm) size bins. The larger background errors of those species allowed the field to be better adjusted, which 10 was crucial for the aerosol analyses in this study. 11

12 **2.3 Experimental design**

We conducted two sets of experiments (NO DA and CONC DA) for January 2015, 2016 and 2017. In 13 both cases, the MEIC_2010 emission inventory was used. The NO_DA experiment initialized a new 14 WRF/Chem forecast every 6 hr starting at 00 UTC on 20 December of the previous year to spin up the aerosol 15 fields and was run through 23 UTC on 31 January. Only the simulations in January were used for the analysis. 16 In the NO DA experiment, the chemical/aerosol fields were simply carried over from cycle to cycle (similar 17 to a continuous aerosol forecast), while the meteorological IC/BC were updated from GFS analysis data every 18 6 hr to prevent the meteorological simulation from drifting. For CONC_DA, the GSI 3DVAR system updated 19 the MOSAIC aerosol variables every hour starting from 00 UTC on 1 January. The background of the first 20 cycle at 00 UTC on 1 January was obtained from the NO DA experiment, and all subsequent cycle were 21 derived from the previous cycle's 1-hr forecast. In CONC_DA, the GFS analysis data were interpolated from 22 a 6-hr frequency to a 1-hr frequency and were then used to update the meteorological IC/BC in each 1-hr cycle. 23 The newly added heterogeneous reactions were activated in both sets of experiments. 24

1 2.4 Distinguishing the impacts of meteorological conditions and emissions

As introduced in section 1, interannual air quality changes are strongly influenced by both emissions and 2 meteorological conditions. It is challenging to distinguish and quantify the impacts of these two aspects solely 3 based on observations or modeling. In our case, the impacts of meteorological conditions are diagnosed by 4 analyzing the differences between two sets of modeling simulations (with the same emission inventory but 5 different meteorology conditions). For NO DA, the emission inputs for January of the three years (2015-2017) 6 were all from the MEIC_2010 emission inventory, and the only differences among the simulations of these 7 three months were the meteorological conditions, which were acquired from the GFS 6-hr analysis data. 8 Therefore, we can assume that the differences in the simulated NO DA PM_{2.5} concentrations among the three 9 months were driven purely by differences in the meteorological conditions (similar to Xu et al. 2017). 10 However, it is difficult to distinguish the impacts of emissions by using the same approach. As temporary 11 emission control measures were applied according to the pollution severity (alarm level), the emission 12 reduction ratios actually continued to change during the winter season, and thus, no exact emission reduction 13 ratios were provided for those days. Nevertheless, the simulation approach with different emission scenarios 14 is simply impossible when lacking exact emission reduction ratios. Instead, we subtracted the meteorological 15 effects from the total effects by utilizing the reanalysis data and pure model simulations. The CONC_DA 16 result, in which the hourly surface PM_{2.5} observations from 531 groups of sites were utilized, can be treated 17 as a reanalysis dataset that reflects the actual conditions (very close to the observations). Therefore, the 18 differences in the assimilated CONC DA PM_{2.5} concentrations among the three months reflect the combined 19 effects of both meteorological conditions and emissions. As the two experiments were generated on gridded 20 aerosol fields, we can separate the effects of emissions from the collective effect by subtracting the NO DA 21 differences from the CONC_DA differences. Hence, we can better comprehend how meteorological 22 conditions and emissions play different roles in driving the changes among the three years. Table 2 illustrates 23 this approach by taking 2015 and 2016 as an example. However, some uncertainties might be associated with 24 this approach, as will be discussed in detail in section 4.2. 25

Table 2. The approach used to distinguish the different impacts of meteorological conditions and emissions by calculating them from different scenarios (taking 2015 and 2016 as an example).

A. Assimilated total changes	CONC_DA_2016- CONC_DA_2015	Reflecting the combined effect of all driving factors, e.g., emissions and meteorological conditions, from 2015 to 2016
B. Simulated changes due to meteorological differences	NO_DA_2016- NO_DA_2015	As NO_DA_2015 and NO_DA_2016 were conducted with same emissions but different meteorological conditions, the differences reflect the effects due to meteorological differences from 2015 to 2016
C. Calculated changes due to emission differences = (A-B)	(CONC_DA_2016- CONC_DA_2015) - (NO_DA_2016- NO_DA_2015)	Mostly reflecting the effects from emission differences between 2015 and 2016

3 **3.** Evaluation of the assimilated PM_{2.5}

1 2

This section presents the results from the NO_DA and assimilation experiments outlined above. In slight
contrast to S12 and Jiang13, our purpose was to reproduce the spatial-temporal variations in the surface PM_{2.5}
within the reanalysis dataset rather than to provide the IC of aerosol fields for improving forecasts.

Figure 3 shows the observed and modeled monthly averages of the surface PM_{2.5} for January 2015, 2016 7 and 2017. Nine regions are illustrated as rectangles in the figure: North China Plain (NEC), northeastern China 8 (NEC), Energy Golden Triangle (EGT), Xinjiang (XJ), Fenwei Plain (FWP), Sichuan Basin (SB), Central 9 China (CC), Yangtze River Delta (YRD), and Pearl River Delta (YRD). Both the observations and the model 10 show that high values are mostly observed in NCP, FWP, SB and CC. In the NO_DA case, the model results 11 are overpredicted in SB, NCP and CC for all three months, while the overestimations are more severe in SB. 12 The NO DA case generally overestimates (underestimates) the surface PM_{2.5} in NCP, SB and CC (XJ and 13 FWP) in the three years, potentially indicating that the 2010 emissions are not appropriate for the 2015-2017 14 simulations with overestimations (underestimations). As discussed in section 2.1, the large area of 15 overestimation is consistent with the national reductions in SO_2 , NO_x and $PM_{2.5}$ anthropogenic emissions 16 (Zheng et al., 2018); however, the underestimations in XJ and FWP also indicate the introduction of new 17 emission sources to these two regions. 18

1	Compared to the NO_DA case, the CONC_DA assimilation experiment effectively reproduces the spatial
2	distribution of surface $PM_{2.5}$ for the three months in terms of the relatively higher values observed in NCP, SB
3	and CC and in some "hot spots" (NEC, FWP, and XJ), which are closer to the observations.
4	Basic statistical measures, including the bias (BIAS), standard deviation (STDV), root-mean-square error
5	(RMSE) and correlation coefficient (CORR), were applied to evaluate the experiments. Figure 4 shows the
6	time series of the BIAS, STDV and RMSE for all the data used in the entire domain. The statistics were
7	calculated for each 1-hr DA cycle. After quality control, the number of $PM_{2.5}$ observations used in the DA
8	process differed; the number of observations was normally approximately 500-520 but reached a minimum of
9	320-450 occasionally due to the data availability. From the time series, we can see that the BIAS, STDV and
10	RMSE are greatly improved in the CONC_DA case. The maximum BIAS values are approximately 50 μ g m ⁻
11	3 for January 2015 and approximately 80 μg m $^{-3}$ for 2016 and 2017 in NO_DA, while they are reduced to

approximately $\pm 5 \ \mu g \ m^{-3}$ in CONC_DA. The STDV and RMSE are also reduced by at least 50% most of

13 the time.

12

Figure 5 shows the spatial distributions of the error statistics (BIAS, RMSE and CORR) at each 14 observational site (with more than 2/3 valid data in the month) in January 2015, 2016 and 2017. We start with 15 2015 and then address the differences with comparisons in 2016 and 2017. In 2015 in the NO_DA case, the 16 surface PM_{2.5} concentrations are generally overestimated by 20-60 µg m⁻³ in eastern China (NCP, SB, CC, 17 PRD and YRD) but are underestimated in NEC, FWP, EGT and especially XJ. The high/low BIAS values in 18 eastern/western China are greatly corrected in CONC_DA. Consistent with the BIAS changes in CONC_DA, 19 the RMSE and CORR distributions in eastern China and NEC are also greatly improved; the RMSE is reduced 20 by at least 50%, and the CORR increases to almost above 0.8-0.9. The inhomogeneous distributions of the 21 BIAS in NO_DA in 2016 and 2017 are very similar to that in 2015 (overestimated in eastern China but 22 underestimated in NEC, EGT and XJ). However, the high biases in CC and PRD and the low biases in XJ are 23 even larger in 2016 and 2017. Similar to the comparisons between NO_DA and CONC_DA for the year 2015, 24 improvements are generally achieved for almost all the regions in both 2016 and 2017. The statistics for the 9 25

1 regions are listed in Table 3.

2

3

Statistics	Sites	Pairs of	В	IAS	R	MSE	CORR		
Statistics	51105	data	NO_DA	CONC_DA	NO_DA	CONC_DA	NO_DA	CONC_DA	
2015									
NCP	67	46699	19.38	2.08	68.09	24.26	0.72	0.96	
NEC	30	20910	-11.94	-1.04	49.47	21.11	0.59	0.93	
EGT	28	19516	-40.43	5.28	60.62	19.45	0.37	0.90	
XJ	19	13243	-53.76	4.16	71.69	19.74	0.40	0.94	
FWP	27	18819	4.05	1.75	56.71	23.05	0.63	0.93	
SB	48	33456	98.02	0.61	125.76	20.76	0.55	0.94	
CC	49	34153	46.94	-0.38	81.31	21.18	0.46	0.93	
YRD	34	23698	32.22	-0.43	59.90	15.14	0.73	0.96	
PRD	20	13940	19.36	-0.03	47.81	9.10	0.24	0.95	
2016									
NCP	67	46699	20.90	1.41	57.77	20.74	0.78	0.96	
NEC	30	20910	-11.05	0.04	40.91	16.08	0.57	0.94	
EGT	28	19516	-22.55	0.69	39.63	13.75	0.42	0.90	
XJ	19	13243	-72.92	0.25	98.19	27.16	0.51	0.96	
FWP	27	18819	-3.51	1.51	62.04	26.01	0.76	0.94	
SB	48	33456	134.63	2.77	165.38	15.49	0.51	0.92	
CC	49	34153	86.28	1.89	109.09	18.76	0.46	0.92	
YRD	34	23698	46.13	1.03	62.11	13.40	0.73	0.95	
PRD	20	13940	59.79	2.05	74.76	6.51	0.04	0.91	
2017				-		·			
NCP	67	46699	25.75	2.35	82.31	28.91	0.74	0.95	
NEC	30	20910	-11.38	0.01	53.38	21.35	0.64	0.94	
EGT	28	19516	-26.88	1.40	48.83	16.96	0.41	0.90	
XJ	19	13243	-95.92	3.82	125.09	35.65	0.51	0.96	
FWP	27	18819	-6.78	-1.02	89.26	31.69	0.65	0.94	
SB	48	33456	122.82	2.33	149.08	20.08	0.56	0.93	
CC	49	34153	101.22	3.49	132.97	19.50	0.23	0.92	
YRD	34	23698	59.31	2.40	78.02	12.32	0.63	0.93	
PRD	20	13940	35.01	0.04	61.84	9.55	-0.16	0.94	

Table 3. Statistics of the observed and model-simulated surface $PM_{2.5}$ for January 2015, 2016 and 2017 in 9 regions (units are $\mu g m^{-3}$ for BIAS and RMSE).

4

5

6 4. Interannual changes during 2015 through 2017

Given reliable PM_{2.5} reanalysis fields produced by assimilating surface PM_{2.5} (CONC_DA), the change
trends among the three years can be analyzed for not only scattered observational sites but also different

regions. To distinguish the roles of meteorological conditions and emissions in driving these changes, an analysis based on the NO_DA and CONC_DA simulations is performed. As assumed in section 2.4, meteorology-driven changes can be analyzed in the NO_DA simulations with different meteorological conditions but the same emission inventory for different years; however, the changes in the reanalysis data among different years are actually the combination of all the driving forces, including meteorological conditions and emissions. By analyzing both sets of simulations, we can attempt to distinguish the roles of meteorology and emissions in determining these changes.

8 4.1 Spatial distribution

The monthly mean PM_{2.5} differences for January in the three years (2015-2017) are shown in Fig. 6 in 9 terms of the surface concentrations measured at observational sites (Fig. 6a) and those from assimilation 10 experiments (Fig. 6b). The surface observations are mostly reduced from 2015 to 2016 except for a few sites 11 in the southern parts of NCP and FWP and in XJ. For the changes from 2016 to 2017, the surface observations 12 increase at almost all the sites, especially the sites in the southern part of NCP; the only exceptions are the 13 sites along the coastline in YRD. The assimilated (CONC_DA) differences are consistent with the surface 14 observations insomuch that the decreasing trend from 2015 to 2016 and the increasing trend from 2016 to 15 2017 for most of the regions are reproduced. However, for the changes in Tibet, EGT and XJ, where 16 observational sites are sparse, some "cold spots" were artificially generated by CONC_DA due to the scarcity 17 of data and the horizontal length scale set in the assimilation. As already shown in Fig. 3 and indicated here 18 again, January 2016 is the cleanest month among the three years. 19

20

21 4.2 The roles of meteorological conditions and emissions

The surface PM_{2.5} concentrations from both the observations and the assimilation experiments show a decreasing trend from 2015 to 2016 but an increasing trend from 2016 to 2017 for most of the regions in eastern China (Fig. 6). The Chinese government has implemented a strict emission control strategy since 2013, especially in northern China, and thus, emission reductions are expected for each year following 2013. The ambient response from 2015-2017 is contradictory if considering only the reductions in emissions and omitting the changes in meteorological conditions. There are two possible assumptions: the first is that the emission reduction target was not achieved from 2016 to 2017, and the second is that other factors in addition to emissions played more important roles.

The NO_DA differences among the different years are shown in Fig. 6c, which reflects the effect of 5 meteorological condition changes (section 2.4). The effect due to emissions (the other major factor in addition 6 to meteorological conditions) is given by subtracting the NO DA differences from the CONC DA differences 7 (Fig. 6d). We can clearly see that the meteorology played two different roles from 2016 to 2017. It caused a 8 decrease in the ambient concentrations for northern China (NCP and NEC) from 2015 to 2016 but induced a 9 large increase for northern and central China (CC) from 2016 to 2017. This indicates that the meteorological 10 conditions might have differed from 2016 to 2017. After considering the impacts of meteorological conditions, 11 those of emission reductions are still confirmed for these two regions from 2016 to 2017. The contributions 12 from both meteorological conditions and emissions in the 9 regions (defined in Fig. 3) were calculated, and 13 the results are listed in Table 4. The calculations show a reduction of approximately 15-20 µg m⁻³ in PM_{2.5} for 14 the month of January from 2015 to 2016 in northern China (NCP and NEC), but the meteorology played a 15 dominant role (contributing a reduction of approximately 12-21 µg m⁻³ in PM_{2.5}). The changes from 2016 to 16 2017 in NCP and NEC are completely different; meteorological conditions caused an increase in PM_{2.5} of 17 approximately 12-23 µg m⁻³, and emission control measures caused a decrease of 1-8 µg m⁻³ in PM_{2.5}, while the 18 combined effects still showed a PM_{2.5} increase for that region. It is reasonable to say that emissions were 19 indeed reduced for the northern regions from 2016 to 2017. However, the meteorology played an important 20 role in offsetting those emission reductions and leading to an increase in surface concentrations in 2017. 21

Table 4. Modeled ambient PM_{2.5} concentration changes for 2016-2015, 2017-2016 and 2017-2015 in 9
 regions and the contributions of the meteorology (MET) and emissions (EMIS) calculated according to
 Table 2. Units: μg m⁻³.

2010-2013 2017-2010 2017-2013

	Total	MET	EMIS	Total	MET	EMIS	Total	MET	EMIS
NCP	-15.23	-12.52	-2.71	+14.91	+23.16	-8.25	-0.31	+10.65	-10.96
NEC	-20.09	-21.23	+1.14	+11.44	+12.61	-1.18	-8.66	-8.62	-0.04
EGT	-21.69	1.68	-23.37	+4.86	+3.81	+1.05	-16.83	+5.48	-22.31
XJ	+3.69	+0.07	+3.63	+1.85	+0.28	+1.57	+5.54	+0.34	+5.20
FWP	-7.05	-10.19	+3.13	+22.95	+25.62	-2.66	+15.90	+15.43	+0.47
SB	-18.75	+8.72	-27.48	+10.31	+4.02	+6.29	-8.45	+12.74	-21.19
CC	-21.80	+14.73	-36.54	+9.35	+19.36	-10.01	-12.45	+34.09	-46.54
YRD	-10.43	-3.03	-7.40	-11.45	-2.93	-8.52	-21.88	-5.96	-15.92
PRD	-23.48	13.02	-36.50	+12.71	-6.12	+18.83	-10.77	+6.90	-17.67

1	It is worth noting that there are uncertainties in the simulation/assimilation processes. There are three
2	sources of uncertainties in the NO_DA simulation. First, the emission inventories in the NO_DA simulations
3	are obviously not accurate, which may introduce uncertainties into the analysis. Although the basic assumption
4	required only that the emissions stay the same throughout the three years, emission inventory uncertainty-
5	induced errors would be offset in the subtraction process when calculating the year-to-year differences.
6	However it did generate uncertainties. For example, the emissions in SB, CC and PRD were generally
7	overestimated (Fig. 3), which means that the variations in the ambient concentration might have been
8	artificially amplified considering the meteorology impacts (Fig. 6c). In contrast, the emissions in XJ and FWP
9	were underestimated (Fig. 3), and thus, the changes in the ambient concentrations due to meteorological
10	conditions in these two regions might have diminished. From this point of view, if the fixed emissions are
11	more accurate in those years, the results would be more reliable. In the case where "real" emissions are not
12	available and the purpose is to evaluate the contribution of those emissions, uncertainties will be unavoidable
13	and should be emphasized carefully. Second, the meteorological IC/BC conditions in the NO_DA simulations,
14	which were obtained from GFS 6-hr analysis data, also have uncertainties. The biases in meteorological
15	conditions might lead to uncertainties in the PM _{2.5} analysis. Third, the deficiencies associated with the
16	chemistry in the model also generate uncertainties, including missing reactions and the inaccurate
17	parameterization of reactions. These three aspects all originate from the imperfections of current forward
18	models. From another perspective, the accuracy of the CONC_DA assimilation experiment also affects the
19	analysis. For example, the assimilation artificially made some "code spots" in Tibet, EGT and XJ, where
20	observational sites are sparse; this could also induce biases. Finally, the contribution of aerosol-meteorology

feedback was not considered in our calculations. As noted by Gao *et al.* (2017), reduced aerosol feedbacks
due to emission reductions accounted for approximately 10.9% of the total decrease in PM_{2.5} concentrations
in urban Beijing in their APEC study. In our current approach, this effect is integrated into the emissions in
the subtracting process.

5 4.3 Meteorological changes in 2016 and 2017

It is interesting to see that meteorology played different roles in each of the three years. Here, we 6 compared some meteorological parameters to explain the impacts of the meteorology. Differences in the 7 monthly mean planetary boundary layer height (PBLH), surface pressure (PSFC), 2-meter temperature (T2), 8 2-meter relative humidity (RH2) and 10-meter wind speed in different years are given in Fig. 7. The statistics 9 of the differences in these parameters in the 9 regions are listed in Table 5, which shows that the changes in 10 the PSFC and T2 for the periods 2015-2016 and 2016-2017 are different over the whole region. Comparing 11 the parameters between 2015 and 2016, the pressure system is stronger, the temperature is lower, and the wind 12 speed is larger in most regions in the latter; these conditions are favorable for the dispersion of pollution. 13 However, there are some unfavorable conditions, including a lower PBLH and a higher RH (and thus, more 14 heterogeneous reactions with the high RH) in northern and southern China, which may offset the impacts of 15 high pressure systems and low temperatures. Therefore, the combined impacts of these meteorological 16 parameters caused a decrease in the ambient concentration in northern China and an increase in southern China 17 from 2015 to 2016, as shown in Fig. 6. The meteorological changes are different from 2016 to 2017 with a 18 weaker pressure system, higher temperature, smaller wind speed, and lower PBLH in most regions, which 19 caused the pollution to accumulate. As suggested by recent studies, climate change has had important impacts 20 on extreme haze events in northern China based on historical statistical approaches or climate models. Those 21 studies (e.g., Li et al., 2015, Zuo et al., 2015) revealed that wintertime fog-haze days across central and eastern 22 China have a close relationship with the East Asian winter monsoon; in addition, significant weakening 23 (strengthening) of the Siberian high and East Asian trough are the two main factors for extreme cold events 24 and extreme warm events throughout China in winter, while warmth boosts air pollution. Consistent with our 25

study, Zhao *et al.* (2018) noted that a stronger Siberian high period in January 2016 produced a significant decrease in PM_{2.5} concentrations relative to those during weaker periods in other years. The abovementioned studies emphasized that climate change factors and the impacts of emission changes are still difficult to evaluate. Our study used the DA technique in combination with regional models and surface observations to distinguish the impacts of emissions and meteorological conditions to further investigate the year-to-year changes at the regional scale.

Table 5. Statistics of the meteorological differences by region for January 2015, 2016 and 2017.

	PB	LH (me	ter)	P	SFC (Pa)	Т	2 (degre	e)]	RH2 (%))	W	/S10 (m/	s)
	2016	2017	2017	2016	2017	2017	2016	2017	2017	2016	2017	2017	2016	2017	2017
	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
	2015	2016	2015	2015	2016	2015	2015	2016	2015	2015	2016	2015	2015	2016	2015
NCP	27.9	-26.7	1.2	138.5	-30.2	108.4	-4.9	3.3	-1.6	3.0	5.1	8.1	1.15	-0.78	0.37
NEC	22.7	35.3	58.0	117.0	-58.7	58.3	-4.9	4.4	-0.5	-5.7	3.1	-2.6	0.96	-0.38	0.57
EGT	13.6	1.1	14.7	28.0	-8.4	19.7	-4.0	4.0	0.0	10.0	-14.9	-4.9	0.14	-0.50	-0.36
XJ	-0.9	-13.8	-14.7	151.3	-43.1	108.1	-1.3	-0.8	-2.1	5.5	-2.1	3.4	0.36	-0.14	0.22
FWP	67.7	-51.6	16.1	64.6	-12.2	52.4	-3.8	3.4	-0.4	2.8	-0.8	2.0	1.05	-1.00	0.06
SB	9.8	-13.2	-3.4	-15.9	15.9	0.1	-2.4	2.5	0.2	3.9	-1.8	2.0	0.43	-0.02	0.41
CC	34.8	-56.6	-21.9	82.8	-53.2	29.6	-2.5	2.1	-0.3	10.8	0.7	11.5	0.60	-0.07	0.53
YRD	64.7	-22.0	42.7	77.1	-27.8	49.2	-1.7	1.9	0.2	7.8	2.5	10.3	0.89	-0.40	0.49
PRD	-36.1	8.2	-27.9	-16.2	-60.1	-76.3	-0.5	2.4	1.9	11.9	-8.7	3.2	0.94	-0.48	0.46

8 5. Conclusions

To analyze the complex PM_{2.5} pollution in China, the GSI-WRF/Chem aerosol data assimilation system
was updated from the GOCART aerosol scheme to the MOSAIC-4BIN scheme, which is more appropriate
for characterizing anthropogenic emission-relevant aerosol species. Three years (2015-2017) of wintertime
(January) surface PM_{2.5} observations from 1600+ sites were assimilated hourly using the updated 3DVAR
system in the CONC_DA assimilation experiment. A parallel control experiment that did not employ DA
(NO_DA) was also performed.

Both the control and the assimilation experiments were evaluated against the surface PM_{2.5} observations. In the NO_DA experiment, in which the 2010_MEIC emission inventory was used, the modeled PM_{2.5} were severely overestimated in the Sichuan Basin (SB), Central China (CC), Yangtze River Delta (YRD), and Pearl River Delta (PRD) by 98-134, 46-101, 32-59, and 19-60 μ g m⁻³, respectively, which indicated that the emission estimates for 2010 are not appropriate for 2015-2017, as strict emission control strategies were implemented in recent years. Meanwhile, underestimations of 11-12, 53-96, and 22-40 μ g m⁻³ were observed in northeastern China (NEC), Xinjiang (XJ) and the Energy Golden Triangle (EGT), respectively. The assimilation experiment significantly reduced the high biases of surface PM_{2.5} in SB, CC, YRD, and PRD and the low biases in NEC and XJ with biases within $\pm 5 \ \mu$ g m⁻³.

Both the observation and the assimilation experiments showed decreasing ambient concentrations from 7 2015 to 2016 but increasing concentrations from 2016 to 2017 for most of the regions. To distinguish the 8 important factors driving these changes, the reanalysis data from the assimilation experiment and the modeling 9 results from the control experiment were analyzed. The results showed a reduction in PM_{2.5} of approximately 10 15-20 µg m⁻³ for the month of January from 2015 to 2016 in northern China (NCP and NEC), but meteorology 11 played the dominant role (contributing approximately 12-21 μ g m⁻³ of the PM_{2.5} reduction). The changes from 12 2016 to 2017 in NCP and NEC were different; meteorological conditions caused an increase in PM2.5 of 13 approximately 12-23 µg m⁻³, while emission control measures caused a decrease of 1-8 µg m⁻³, and the 14 combined effects still showed a PM_{2.5} increase for that region. The analysis confirmed that meteorology played 15 different roles in 2016 and 2017: the higher pressure system, lower temperatures and higher PBLH in 2016 16 (compared with 2015) were favorable for pollution dispersion, whereas the situation was almost the opposite 17 in 2017 (compared with 2016) and led to an increased PM_{2.5} from 2016 to 2017, although emission control 18 strategies were implemented in both years. After considering the impacts of the meteorology, the analysis 19 showed that emissions were indeed reduced from 2015 to 2016 and 2017, especially in NCP for the year 2017 20 (although the surface concentrations increased that year). The analysis also showed that emissions increased 21 in XJ and FWP. 22

There are still large uncertainties in this approach, such as the deficiencies of forward models (including inaccurate emission inputs, uncertainties in the meteorological IC/BC, and the chemistry mechanism) and the assimilation process, and the imperfection of the aerosol-meteorology feedbacks in the model simulation 1 generated large biases in the analysis. The most straightforward approach is thus to directly estimate the

- 2 emissions by data assimilation, which will be the topic of a separate study.
- 3

4 Acknowledgement

5 This work was supported by the National Key R&D Program on Monitoring, Early Warning and 6 Prevention of Major Natural Disasters under grant (2017YFC1501406), the National Natural Science 7 Foundation of China (Grant No. 41807312) and Basic R&D special fund for central scientific research 8 institutes (IUMKYSZHJ201701). NCAR is sponsored by US National Science Foundation.

9 **References**

Chen, D., Liu, Z. Q., Fast, J., and Ban, J. M.: Simulations of sulfate-nitrate-ammonium (SNA) aerosols during 10 the extreme haze events over northern China in October 2014, Atmos. Chem. Phys., 16, 10707-10724, 10.5194/acp-11 16-10707-2016, 2016. 12 Chen, F., and Dudhia, J.: Coupling an advanced land surface-hydrology model with the Penn State-NCAR 13 MM5 modeling system. Part I: Model implementation and sensitivity, Mon. Weather Rev., 129, 569-585, 2001. 14 Cheng, Y., Zheng, G., Wei, C., Mu, O., Zheng, B., Wang, Z., Gao, M., Zhang, O., He, K., Carmichael, G., 15 Poschl, U., and Su, H.: Reactive nitrogen chemistry in aerosol water as a source of sulfate during haze events in 16 China, Sci. Adv., 2, e1601530, doi:10.1126/sciadv.1601530, 2016. 17 China National Environmental Monitoring Center, Ambient Air Quality Monthly Report 2015-01/2016-18 01/2017-01; http://www.cnemc.cn/kqzlzkbgyb2092938.jhtml. 19 China National Environmental Monitoring Center, PM2.5 Auto-Monitoring Instrument Technical Standard 20 and Requirement, Beijing, 2013. 21 Chin, M., Savoie, D. L., Huebert, B. J., Bandy, A. R., Thornton, D. C., Bates, T. S., Quinn, P. K., Saltzman, E. 22 S., and De Bruyn, W. J.: Atmospheric sulfur cycle simulated in the global model GOCART: Comparison with field 23 observations and regional budgets, J. Geophys. Res.-Atmos., 105, 24689-24712, doi:10.1029/2000jd900385, 2000. 24 25 Chin, M., Ginoux, P., Kinne, S., Torres, O., Holben, B. N., Duncan, B. N., Martin, R. V., Logan, J. A., Higurashi, A., and Nakajima, T.: Tropospheric aerosol optical thickness from the GOCART model and comparisons 26 with satellite and Sun photometer measurements, J. Atmos. Sci., 59, 461-483, doi:10.1175/1520-27 0469(2002)059<0461:Taotft>2.0.Co;2, 2002. 28 Chou, M.-D., and M.J. Suarez.: An efficient thermal infrared radiation parameterization for use in general 29 30 circulation models, NASA Tech. Memo., TM 104606, vol. 3, 25 pp., NASA Goddard Space Flight Cent., Greenbelt, Md. 1994 31 Collins, W. D., Rasch, P. J., Eaton, B. E., Khattatov, B. V., Lamarque, J. F., and Zender, C. S.: Simulating 32 aerosols using a chemical transport model with assimilation of satellite aerosol retrievals: Methodology for 33 34 INDOEX, J. Geophys. Res.-Atmos., 106, 7313-7336, doi:10.1029/2000jd900507, 2001.

35 Elbern, H., Strunk, A., Schmidt, H., and Talagrand, O.: Emission rate and chemical state estimation by 4-

dimensional variational inversion, Atmos Chem Phys, 7, 3749-3769, DOI 10.5194/acp-7-3749-2007, 2007.

Fast, J. D., Gustafson, W. I., Easter, R. C., Zaveri, R. A., Barnard, J. C., Chapman, E. G., Grell, G. A., and Peckham, S. E.: Evolution of ozone, particulates, and aerosol direct radiative forcing in the vicinity of Houston using a fully coupled meteorology-chemistry-aerosol model, J. Geophys. Res.-Atmos., 111, D21305, doi:10.1029/2005jd006721, 2006.

Gao, M., Liu, Z., Wang, Y., Lu, X., Ji, D., Wang, L., Li, M., Wang, Z., Zhang, Q., Carmichael., G. R.: Distinguishing the roles of meteorology, emission control measures, regional transport, and co-benefits of reduced aerosol feedbacks in "APEC Blue". Atmos. Environ., 167, 476-486, doi:10.1016/j.atmosenv.2017.08.054, 2017.

Grell, G. A., Peckham, S. E., Schmitz, R., McKeen, S. A., Frost, G., Skamarock, W. C., and Eder, B.: Fully coupled "online" chemistry within the WRF model, Atmos. Environ., 39, 6957-6975, doi:10.1016/j.atmosenv.2005.04.027, 2005.

Grell, G. A., and Devenyi, D.: A generalized approach to parameterizing convection combining ensemble and data assimilation techniques, Geophys. Res. Lett., 29, Artn 1693.Doi 10.1029/2002gl015311, 2002.

Han, B., Zhang, R., Yang, W., Bai, Z., Ma, Z., and Zhang, W.: Heavy air pollution episodes in Beijing during January 2013: inorganic ion chemistry and source analysis using Highly Time-Resolved Measurements in an urban site, Atmos. Chem. Phys. Discuss., 15, 11111-11141, doi:10.5194/acpd-15-11111-2015, 2015.

He, K.B., Multi-resolution emission Inventory for China (MEIC): model framework and 1990-2010 anthropogenic emissions. In: Presented on the international Global Atmospheric Chemistry Conference, September 17-21, Beijing, China 2012

Hong, C. P., Zhang, Q., He, K. B., Guan, D. B., Li, M., Liu, F., and Zheng, B.: Variations of China's emission estimates: response to uncertainties in energy statistics, Atmos. Chem. Phys., 17, 1227-1239, 10.5194/acp-17-1227-2017, 2017.

Hong, S. Y., Noh, Y., and Dudhia, J.: A new vertical diffusion package with an explicit treatment of entrainment processes, Mon. Weather Rev., 134, 2318-2341, 2006.

Jiang, Z., Z. Liu, T. Wang, C. S. Schwartz, H.-C. Lin, and F. Jiang.: Probing into the impact of 3DVAR assimilation of surface PM10 observations over China using process analysis, J. Geophys. Res. Atmos., 118, 6738–6749, doi:10.1002/jgrd.50495, 2013.

Krotkov, N. A., McLinden, C. A., Li, C., Lamsal, L. N., Celarier, E. A., Marchenko, S. V., Swartz, W. H.,
Bucsela, E. J., Joiner, J., Duncan, B. N., Boersma, K. F., Veefkind, J. P., Levelt, P. F., Fioletov, V. E., Dickerson, R.
R., He, H., Lu, Z. F., and Streets, D. G.: Aura OMI observations of regional SO2 and NO2 pollution changes from 2005 to 2015, Atmos Chem Phys, 16, 4605-4629, 10.5194/acp-16-4605-2016, 2016.

Liu, J., Han, Y., Tang, X., Zhu, J., Zhu, T.: Estimating adult mortality attributable to PM2.5 exposure in China with assimilated PM2.5 concentrations based on a ground monitoring network. Sci. Total Environ., 568, 1253-1262, 2016.

Liu, Z. Q., Liu, Q. H., Lin, H. C., Schwartz, C. S., Lee, Y. H., and Wang, T. J.: Three-dimensional variational assimilation of MODIS aerosol optical depth: Implementation and application to a dust storm over East Asia, J. Geophys. Res.-Atmos., 116, Artn D23206. doi:10.1029/2011jd016159, 2011.

Lei, Y., Zhang, Q., He, K. B., and Streets, D. G.: Primary anthropogenic aerosol emission trends for China,
1990-2005, Atmos. Chem. Phys., 11, 931-954, 10.5194/acp-11-931-2011, 2011.

Li, M., Zhang, Q., Streets, D. G., He, K. B., Cheng, Y. F., Emmons, L. K., Huo, H., Kang, S. C., Lu, Z., Shao,
M., Su, H., Yu, X., and Zhang, Y.: Mapping Asian anthropogenic emissions of non-methane volatile organic
compounds to multiple chemical mechanisms, Atmos. Chem. Phys., 14, 5617-5638, 10.5194/acp-14-5617-2014,
2014.

Li, Q., Zhang, R. H., and Wang, Y.: Interannual variation of the wintertime fog-haze days across central and eastern China and its relation with East Asian winter monsoon, Int. J. Climatol., 36, 346-354, 10.1002/joc.4350, 2016.

Li, G., Bei, N., Cao, J., Huang, R., Wu, J., Feng, T., Wang, Y., Liu, S., Zhang, Q., Tie, X., and Molina, L. T.: A possible pathway for rapid growth of sulfate during haze days in China, Atmos. Chem. Phys., 17, 3301-3316, doi:10.5194/acp-17-3301-2017, 2017.

McKeen, S. A., Wotawa, G., Parrish, D. D., Holloway, J. S., Buhr, M. P., Hubler, G., Fehsenfeld, F. C., and Meagher, J. F.: Ozone production from Canadian wildfires during June and July of 1995, J. Geophys. Res.-Atmos., 107, 4192, doi:10.1029/2001jd000697, 2002.

Mlawer, E. J., Taubman, S. J., Brown, P. D., Iacono, M. J., and Clough, S. A.: Radiative transfer for inhomogeneous atmospheres: RRTM, a validated correlated-k model for the longwave, J. Geophys. Res.-Atmos., 102, 16663-16682, 1997.

Moch, J. M., Dovrou, E., Mickley, L. J., Keutsch, F. N., Cheng, Y., Jacob, D. J., Jiang, J. K., Li, M., Munger,
J. W., Qiao, X. H., and Zhang, Q.: Contribution of Hydroxymethane Sulfonate to Ambient Particulate Matter: A
Potential Explanation for High Particulate Sulfur During Severe Winter Haze in Beijing, Geophys Res Lett, 45, 11969-11979, 10.1029/2018gl079309, 2018.

Pagowski, M., Grell, G. A., McKeen, S. A., Peckham, S. E., and Devenyi D.: Three dimensional variational data assimilation of ozone and fine particulate matter observations: some results using the Weather Research and Forecasting-Chemistry model and Grid-point Statistical Interpolation, Q. J. R. Meteorol. Soc., 136(653), 2010.

Sun, Y. L., Wang, Z. F., Du, W., Zhang, Q., Wang, Q. Q., Fu, P. Q., Pan, X. L., Li, J., Jayne, J., and Worsnop,
D. R.: Long-term real-time measurements of aerosol particle composition in Beijing, China: seasonal variations, meteorological effects, and source analysis, Atmos. Chem. Phys., 15, 10149-10165, 10.5194/acp-15-10149-2015, 2015.

Schwartz, C. S., Z. Liu, H. C. Lin, and S. A. McKeen.: Simultaneous three-dimensional variational assimilation of surface fine particulate matter and MODIS aerosol optical depth, J. Geophys. Res.-Atmos., 117, 2012.

Shao, J., Chen, Q., Wang, Y., Lu, X., He, P., Sun, Y., Shah, V., Martin, R. V., Philip, S., Song, S., Zhao, Y., Xie,
Z., Zhang, L., and Alexander, B.: Heterogeneous sulfate aerosol formation mechanisms during wintertime Chinese haze events: Air quality model assessment using observations of sulfate oxygen isotopes in Beijing, Atmos. Chem. Phys. Discuss., https://doi.org/10.5194/acp-2018-1352, in review, 2019.

The Central People's Government of the People Republic of China. Guiding Opinions on Promoting the Joint
 Prevention and Control of Air Pollution to Improve Regional Air Quality, Beijing. 2010.

The Central People's Government of the People Republic of China. Atmospheric Pollution Prevention and
 Control Action Plan. 2013.

Wang, L. T., Wei, Z., Yang, J., Zhang, Y., Zhang, F. F., Su, J., Meng, C. C., and Zhang, Q.: The 2013 severe
haze over southern Hebei, China: model evaluation, source apportionment, and policy implications, Atmos. Chem.
Phys., 14, 3151-3173, 10.5194/acp-14-3151-2014, 2014a.

Wang, W., Maenhaut, W., Yang, W., Liu, X. D., Bai, Z. P., Zhang, T., Claeys, M., Cachier, H., Dong, S. P., and
Wang, Y. L.: One-year aerosol characterization study for PM_{2.5} and PM10 in Beijing, Atmos. Pollut. Res., 5, 554562, 10.5094/Apr.2014.064, 2014b.

Wang, Y. S., Yao, L., Wang, L. L., Liu, Z. R., Ji, D. S., Tang, G. Q., Zhang, J. K., Sun, Y., Hu, B., and Xin, J.
Y.: Mechanism for the formation of the January 2013 heavy haze pollution episode over central and eastern China,
Sci. China Earth Sci., 57, 14-25, 10.1007/s11430-013-4773-4, 2014c.

Wang, G., Zhang, R., Gomez, M. E., Yang, L., Levy Zamora, M., Hu, M., Lin, Y., Peng, J., Guo, S., Meng, J.,
Li, J., Cheng, C., Hu, T., Ren, Y., Wang, Y., Gao, J., Cao, J., An, Z., Zhou, W., Li, G., Wang, J., Tian, P., Marrero-Ortiz, W., Secrest, J., Du, Z., Zheng, J., Shang, D., Zeng, L., Shao, M., Wang, W., Huang, Y., Wang, Y., Zhu, Y., Li,
Y., Hu, J., Pan, B., Cai, L., Cheng, Y., Ji, Y., Zhang, F., Rosenfeld, D., Liss, P. S., Duce, R. A., Kolb, C. E., and
Molina, M. J.: Persistent sulfate formation from London Fog to Chinese haze, Proc. Natl. Acad. Sci., 113, 13630-13635, doi:10.1073/pnas.1616540113, 2016.

Wang, G., Zhang, F., Peng, J., Duan, L., Ji, Y., Marrero-Ortiz, W., Wang, J., Li, J., Wu, C., Cao, C., Wang, Y.,
Zheng, J., Secrest, J., Li, Y., Wang, Y., Li, H., Li, N., and Zhang, R.: Particle acidity and sulfate production during severe haze events in China cannot be reliably inferred by assuming a mixture of inorganic salts, Atmos. Chem.
Phys., 18, 10123-10132, doi:10.5194/acp-18-10123-2018, 2018

Wild, O., Zhu, X., and Prather, M. J.: Fast-j: Accurate simulation of in- and below-cloud photolysis in tropospheric chemical models, J. Atmos. Chem., 37, 245-282, 2000.

Xu, J., Chang, L., Yan, F., and He, J.: Role of climate anomalies on decadal variation in the occurrence of wintertime haze in the Yangtze River Delta, China. Science of the Total Environment,599-600, 918-925, doi:10.1016/j.scitotenv.2017.05.015, 2017.

Yang, Y. R., Liu, X. G., Qu, Y., An, J. L., Jiang, R., Zhang, Y. H., Sun, Y. L., Wu, Z. J., Zhang, F., Xu, W. Q., and Ma, Q. X.: Characteristics and formation mechanism of continuous hazes in China: a case study during the autumn of 2014 in the North China Plain, Atmos. Chem. Phys., 15, 8165-8178, 10.5194/acp-15-8165-2015, 2015.

Zhang, L., Shao, J., Lu, X., Zhao, Y., Hu, Y., Henze, D. K., Liao, H., Gong, S., Zhang, Q.: Sources and Processes Affecting Fine Particulate Matter Pollution over North China: An Adjoint Analysis of the Beijing APEC Period. Environ. Sci. Technol., 50, 8731-8740, 2016.

Zhang, Q., Streets, D. G., Carmichael, G. R., He, K. B., Huo, H., Kannari, A., Klimont, Z., Park, I. S., Reddy, S., Fu, J. S., Chen, D., Duan, L., Lei, Y., Wang, L. T., and Yao, Z. L.: Asian emissions in 2006 for the NASA INTEX-B mission, Atmos. Chem. Phys., 9, 5131-5153, 2009.

Zaveri, R. A., and Peters, L. K.: A new lumped structure photochemical mechanism for large-scale applications, J. Geophys. Res.-Atmos., 104, 30387-30415, 1999.

Zaveri, R. A., Easter, R. C., Fast, J. D., and Peters, L. K.: Model for Simulating Aerosol Interactions and Chemistry (MOSAIC), J. Geophys. Res.-Atmos., 113, D13204, doi:10.1029/2007jd008782, 2008.

Zhao, S., Feng, T., Tie, X., Long, X., Li, G., Cao, J., Zhou, W., and An, Z.: Impact of Climate Change on Siberian High and Wintertime Air Pollution in China in Past Two Decades, Earth's Future, 6, 118–133, https://doi.org/10.1002/2017EF000682, 2018.

Zhao, Y., Zhou, Y. D., Qiu, L. P., and Zhang, J.: Quantifying the uncertainties of China's emission inventory for industrial sources: From national to provincial and city scales, Atmos. Environ., 165, 207-221, 10.1016/j.atmosenv.2017.06.045, 2017.

Zheng, B., Zhang, Q., Zhang, Y., He, K. B., Wang, K., Zheng, G. J., Duan, F. K., Ma, Y. L., and Kimoto, T.:
Heterogeneous chemistry: a mechanism missing in current models to explain secondary inorganic aerosol formation
during the January 2013 haze episode in North China, Atmos. Chem. Phys., 15, 2031-2049, 10.5194/acp-15-20312015, 2015.

Zheng, B., Tong, D., Li, M., Liu, F., Hong, C. P., Geng, G. N., Li, H. Y., Li, X., Peng, L. Q., Qi, J., Yan, L.,
Zhang, Y. X., Zhao, H. Y., Zheng, Y. X., He, K. B., and Zhang, Q.: Trends in China's anthropogenic emissions since
2010 as the consequence of clean air actions, Atmos Chem Phys, 18, 14095-14111, 10.5194/acp-18-14095-2018,
2018.

Zhi, G. R., Zhang, Y. Y., Sun, J. Z., Cheng, M. M., Dang, H. Y., Liu, S. J., Yang, J. C., Zhang, Y. Z., Xue, Z.

G., Li, S. Y., and Meng, F.: Village energy survey reveals missing rural raw coal in northern China: Significance in science and policy, Environ. Pollut., 223, 705-712, 10.1016/j.envpol.2017.02.009, 2017.

Zuo, Z. Y., Zhang, R. H., Huang, Y., Xiao, D., and Guo, D.: Extreme cold and warm events over China in wintertime, Int. J. Climatol., 35, 3568-3581, 10.1002/joc.4229, 2015.

1

2 **Tables and Figures**

3 **Table 1.** WRF/Chem model configuration.

Table 2. The approach used to distinguish the different impacts of meteorological conditions and emissions
 by calculating them from different scenarios (taking 2015 and 2016 as an example).

6 **Table 3.** Statistics of the observed and model-simulated surface $PM_{2.5}$ for January 2015, 2016 and 2017 in 9 7 regions (units are $\mu g m^{-3}$ for BIAS and RMSE).

Table 4. Modeled ambient PM_{2.5} concentration changes for 2016-2015, 2017-2016 and 2017-2015 in 9 regions
 and the contributions of the meteorology (MET) and emissions (EMIS) calculated according to Table 2. Units:
 μg m⁻³.

- **Table 5.** Statistics of the meteorological differences by region for January 2015, 2016 and 2017.
- Figure 1. Domain-averaged standard deviations of the background errors ($\mu g kg^{-1}$) as a function of the height for each aerosol variable in three bins: (a) Bin-01: 0.039-0.1 μ m; (b) Bin-02: 0.1-1.0 μ m; (c) Bin-03: 1.0-2.5 μ m.

Figure 2. Spatial distribution of primary $PM_{2.5}$ (the sum of BC, OC, sulfate, nitrate and other unspecified $PM_{2.5}$ emissions), SO₂, NO_x and NH₃ emissions (units are $\mu g m^{-2} S^{-1}$ for PM_{2.5} and mol km⁻² hr⁻¹ for the other three) used in this study.

Figure 3. Observed and modeled monthly average PM_{2.5} concentrations (unit: μg m⁻³) for January 2015 (left),

- 19 2016 (middle) and 2017 (right). Regions defined in red rectangles are as follows: a-NCP (North China Plain),
- 20 b-NEC (northeastern China), c-EGT (Energy Golden Triangle), d-XJ (Xinjiang), e-SB (Sichuan Basin), f-CC
- 21 (Central China), g-YRD (Yangtze River Delta), and h-PRD (Pearl River Delta).

22 Figure 4. Time series of the statistics between the model simulations and observations. Red lines-

CONC_DA minus observations, blue lines –NO_DA minus observations. Statistics include the number of data pairs for comparison, the MEAN-mean bias, the STDV- standard deviation, and the RMS-root mean square error. Left-2015, middle-2016, right-2017. (units are μ g m⁻³ for MEAN, STDV and RMS).

Figure 5. Spatial distributions of the statistics between the model simulations and observations for January 27 2015. Top: NO_DA vs. observations, bottom: CONC_DA vs. observations. BIAS-model minus observation,

28 RMSE-root mean square error, CORR-correlation coefficient. (units are $\mu g m^{-3}$ for BIAS and RMSE).

Figure 6. Observed and modeled ambient $PM_{2.5}$ concentration changes for January 2016-2015 (left), 2017-2016 (middle) and 2017-2015 (right). (a) Observations, (b) assimilated total changes, (c) modeled changes due to meteorological conditions, (d) calculated changes due to emissions. (Units: $\mu g m^{-3}$)

Figure 7. Modeled meteorological changes for 2016-2015 (left), 2017-2016 (middle) and 2017-2015 (right).
(a) PBLH, (b) PSFC, (c) T2, (d) RH2 and (e) 10-m wind speed.

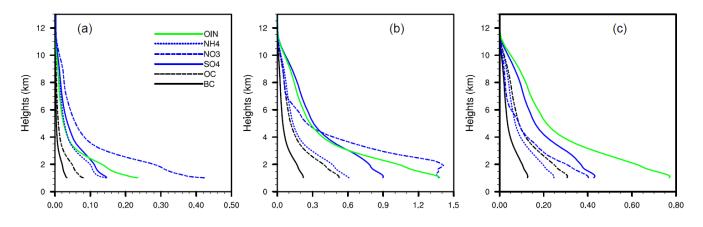


Figure 1. Domain-averaged standard deviations of the background errors (μ g kg⁻¹) as a function of the height for each aerosol variable in three bins: (a) Bin-01: 0.039-0.1 μ m; (b) Bin-02: 0.1-1.0 μ m; (c) Bin-03: 1.0-2.5 μ m.

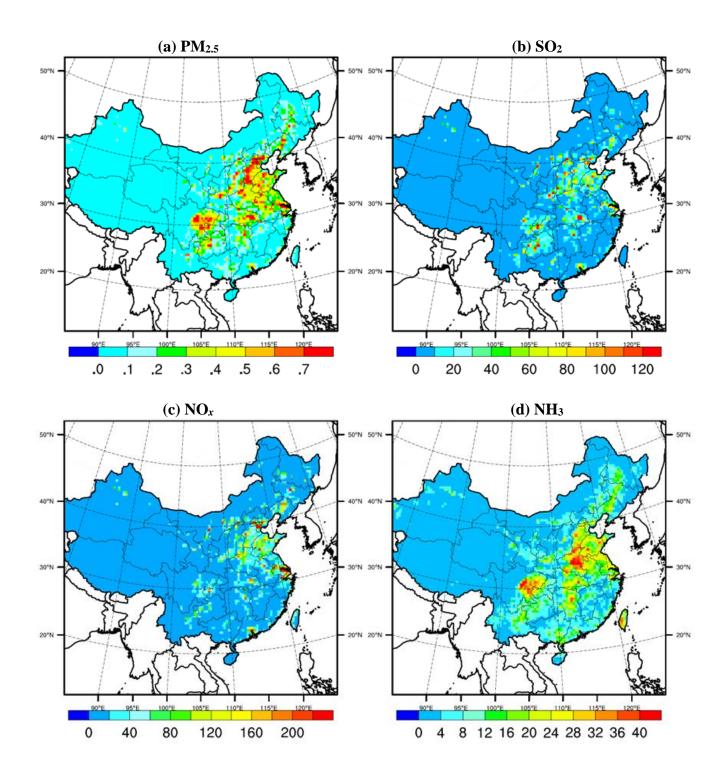


Figure 2. Spatial distribution of primary $PM_{2.5}$ (the sum of BC, OC, sulfate, nitrate and other unspecified $PM_{2.5}$ emissions), SO₂, NO_x and NH₃ emissions (units are $\mu g m^{-2} S^{-1}$ for PM_{2.5} and mol km⁻² hr⁻¹ for the other three) used in this study.

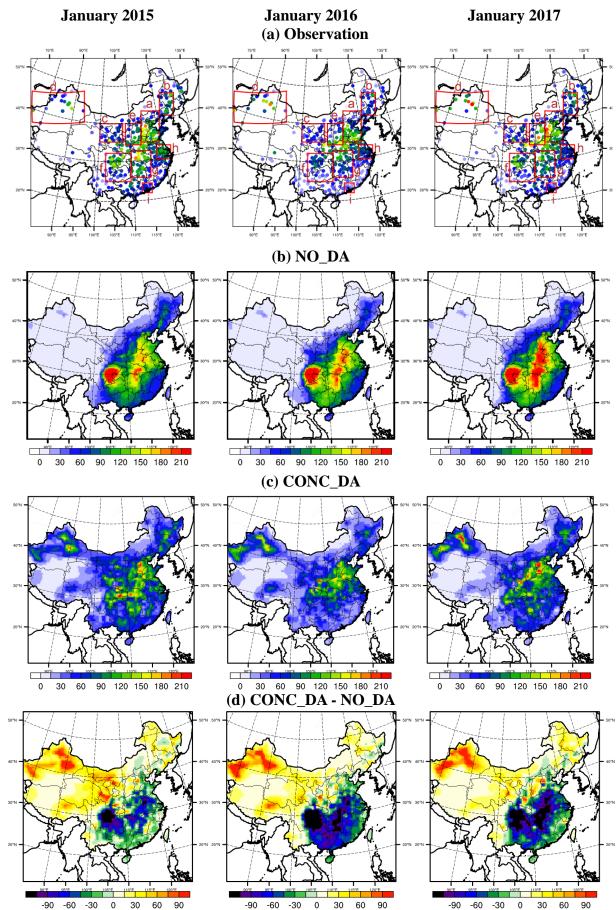


Figure 3. Observed and modeled monthly average PM_{2.5} concentrations (unit: μg m⁻³) for January 2015 (left), 2016 (middle) and 2017 (right). Regions defined in red rectangles are as follows: a-NCP (North China Plain), b-NEC (northeastern China), c-EGT (Energy Golden Triangle), d-XJ (Xinjiang), e-SB (Sichuan Basin), f-CC (Central China), g-YRD (Yangtze River Delta), and h-PRD (Pearl River Delta).

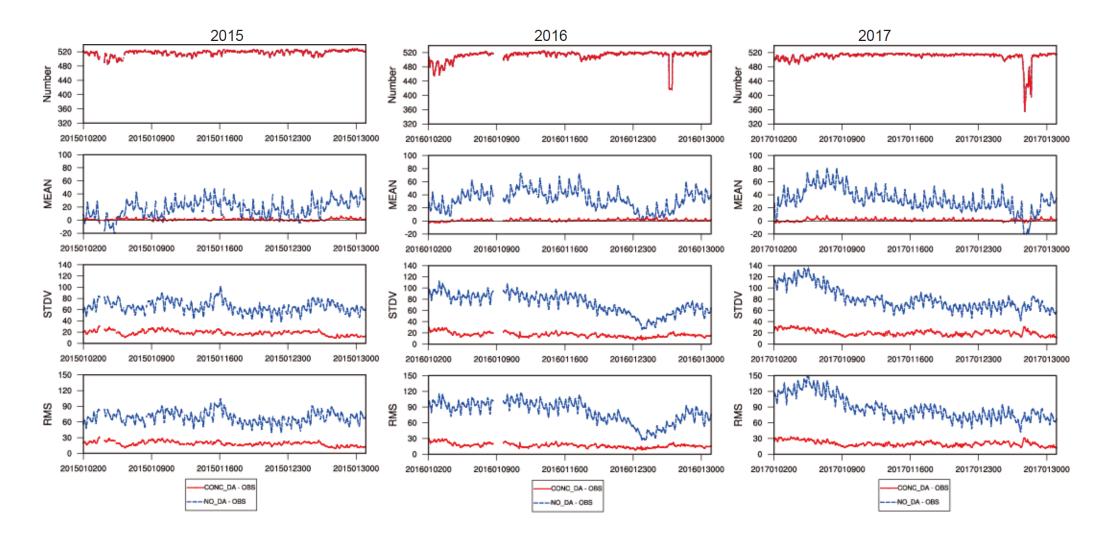
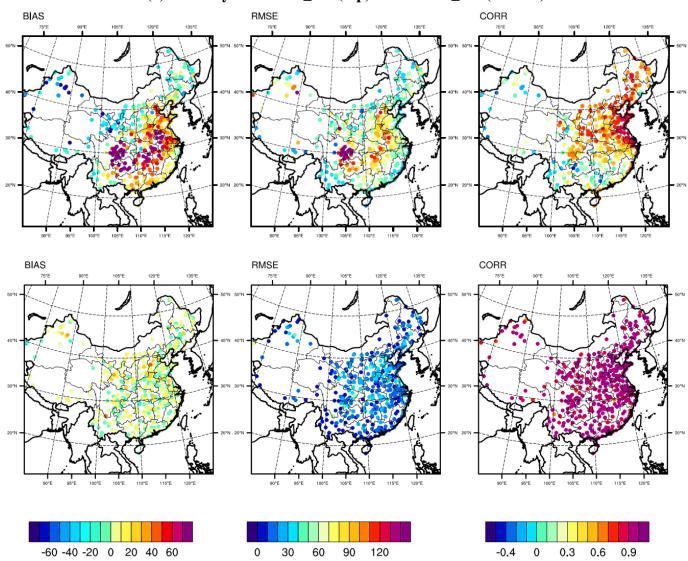


Figure 4. Time series of the statistics between the model simulations and observations. Red lines- CONC_DA minus observations, blue lines –NO_DA minus observations. Statistics include the number of data pairs for comparison, the MEAN-mean bias, the STDV- standard deviation, and the RMS-root mean square error. Left-2015, middle-2016, right-2017. (units are $\mu g m^{-3}$ for MEAN, STDV and RMS).



(a). January 2015 - NO_DA (top) and CONC_DA (bottom)

Figure 5. Spatial distributions of the statistics between the model simulations and observations for January 2015. Top: NO_DA vs. observations, bottom: CONC_DA vs. observations. BIAS-model minus observation, RMSE-root mean square error, CORR-correlation coefficient. (units are $\mu g m^{-3}$ for BIAS and RMSE).

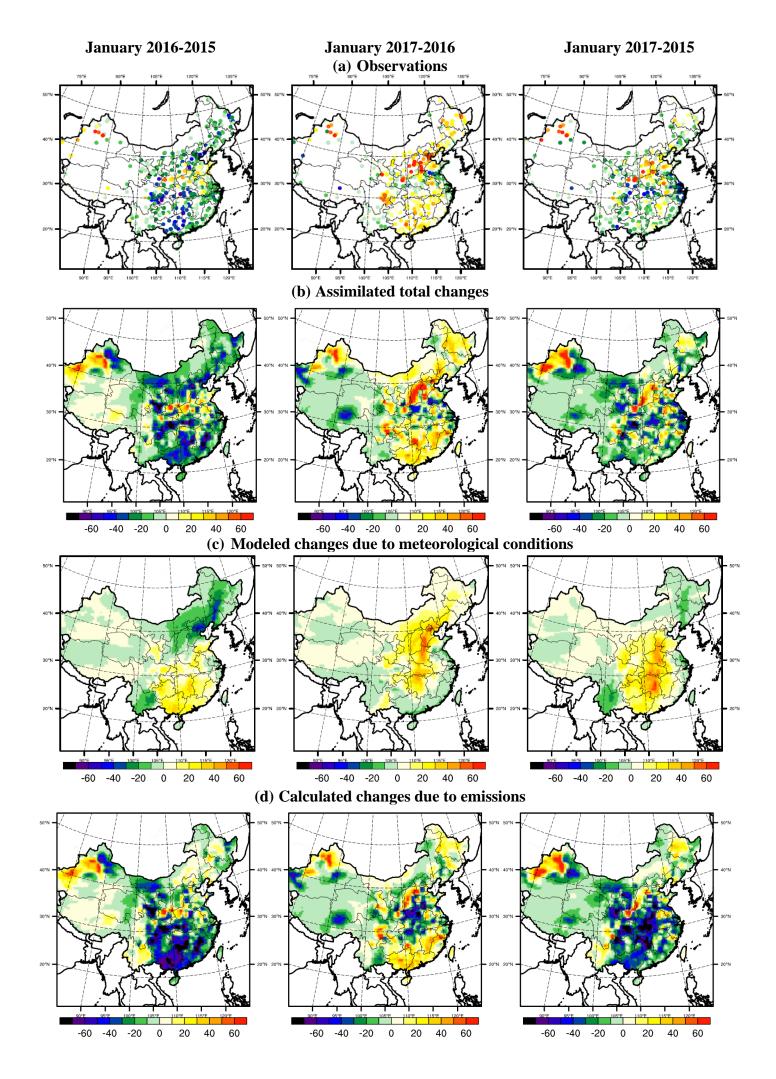


Figure 6. Observed and modeled ambient $PM_{2.5}$ concentration changes for 2016-2015 (left), 2017-2016 (middle) and 2017-2015 (right). (a) Observations, (b) assimilated total changes, (c) modeled changes due to meteorological conditions, (d) calculated changes due to emissions. (Units: $\mu g m^{-3}$)

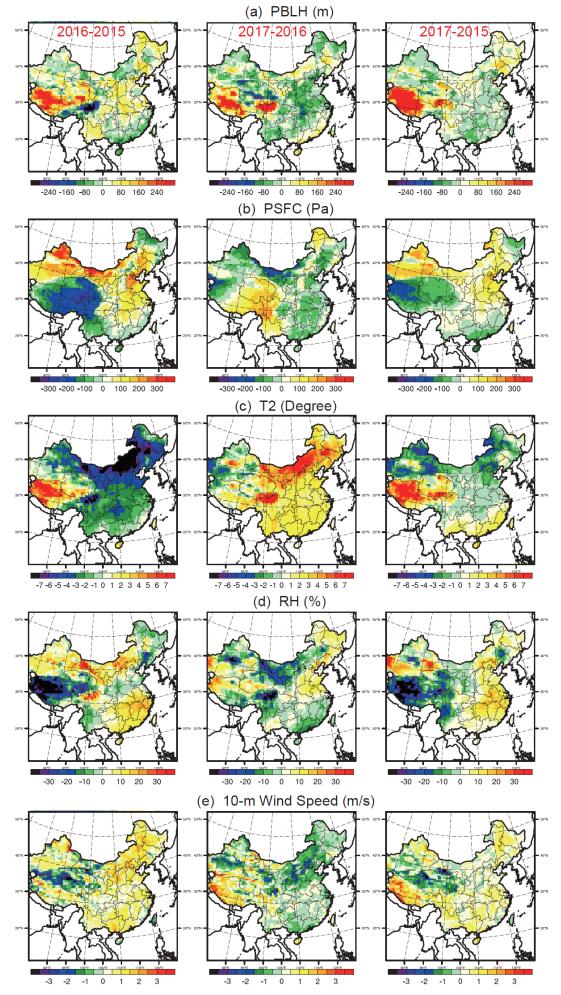


Figure 7. Modeled meteorological changes for 2016-2015 (left), 2017-2016 (middle) and 2017-2015 (right). (a) PBLH, (b) PSFC, (c) T2, (d) RH2 and (e) 10-m wind speed.