



1 Hygroscopic growth effect on aerosol light scattering in the urban area of
2 Beijing: a long-term measurement by a wide-range and high-resolution
3 humidified nephelometer system

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10

11 **Abstract:**

12 Hygroscopicity is an important feature of ambient aerosols, which is very crucial to the
13 study of light extinction, radiation force, and formation mechanism. The light scattering
14 hygroscopic growth factor ($f(\text{RH})$) is an important parameter which is usually measured
15 by the humidified nephelometer system and could better describe the aerosol
16 hygroscopicity under wide particle size range and continuous relative humidity (RH).
17 The $f(\text{RH})$ can be applied to the establishment of a parameterization scheme for light
18 extinction, the calculation of hygroscopicity parameter (κ), and also the estimation of
19 aerosol liquid water content (ALWC). However, the humidified nephelometer system
20 in the previous studies could only observe the $f(\text{RH})$ below 90% due to the larger error
21 of the sensor under high RH (>90%). Furthermore, the $f(\text{RH})$ observations in North
22 China Plain needs to be greatly strengthened both in the temporal resolution and the
23 observation duration. In view of this, an improved high-resolution humidified
24 nephelometer system was established to observe the $f(\text{RH})$ of $\text{PM}_{2.5}$ for a wide RH range
25 between 30%-96% in the urban area of Beijing over three seasons (winter, summer, and
26 autumn) in 2017. It was found that the $f(80\%)$ at 525nm of $\text{PM}_{2.5}$ was evidently higher
27 under the polluted conditions and highly correlated with the fractions of all the water-
28 soluble ions. A two-parameter fit equation was selected to fit the observed $f(\text{RH})$ data.
29 For each season, the fitting curve under the very clean condition was lower than that of
30 other conditions. And the $f(\text{RH})$ points of polluted conditions were more concentrated
31 with higher fitting R^2 for summer and autumn data. The hygroscopicity of aerosol under
32 higher RH was probably enhanced when compared with the data in the previous study
33 conducted in NCP. In summer, the fitting $f(\text{RH})$ showed a significant dependence on
34 wavelength for each pollution condition. However, there was an opposite performance
35 in the $f(\text{RH})$ curves of different wavelengths for the very clean condition in winter. It
36 was shown that The simulation showed that the maximum uncertainty of $f(\text{RH})$ was
37 less than 10%.



38 1. Introduction

39 Atmospheric aerosols influence the atmospheric visibility and earth-atmosphere
40 radiation budget directly by scattering and absorbing solar radiation (Charlson et al.,
41 1992; Schwartz et al., 1995; DeBell et al., 2006), and indirectly by modifying
42 microphysical properties and lifetime of clouds (Twomey, 1974; Albrecht, 1989;
43 Rosenfeld, 2000). The particle size, single scattering albedo (SSA), asymmetry
44 parameter (AP), and refractive index are the key parameters in estimating the aerosol
45 radiative forcing; and these parameters are strongly dependent on the relative humidity
46 (RH). Water uptake can result in enlarged particle sizes, SSA, and AP, change the
47 scattering phase functions, and decreased the refractive index (Cheng et al., 2008).
48 Moreover, numerous studies demonstrate that reaction in aerosol liquid water is an
49 important pathway of secondary aerosol formation, and thus plays a significant role in
50 the overall aerosol chemical composition (McMurry and Wilson, 1983; Ravishankara
51 1997; Kolb et al., 2010; Jang et al., 2002; Liu et al., 2012; Cheng et al., 2016).
52 Furthermore, water uptake is also important in the remote-sensing measurements or
53 satellite retrievals of aerosol optical properties (Wang and Martin, 2007; Zieger et al.,
54 2012).

55 Particle size and chemical composition are the main factors affecting the moisture
56 absorption ability of aerosol particles (Köhler, 1936; Hinds, 2011; Petters and
57 Kreidenweis, 2007; Seinfeld and Pandis, 2016). Particle size mainly affects the extent
58 of the Kelvin effect, while the chemical composition is the most crucial factor. Some
59 simulations demonstrated that the changes of aerosol size distribution could only lead
60 to slight variations in hygroscopic growth factor if the chemical compositions were
61 fixed (Fierz-Schmidhauser et al., 2010; Liu 2015; Kuang et al., 2017).

62 Generally, aerosol hygroscopicity could be described by the diameter growth factor
63 ($g(\text{RH})$), light scattering hygroscopic growth factor ($f(\text{RH})$), and hygroscopicity
64 parameter (κ). Size-resolved $g(\text{RH})$ could be directly measured by the Humidified
65 Tandem Differential Mobility Analyzer (HTDMA) (Liu et al. 1978; Swietlicki et al.,
66 2008). Due to the limitation of the HTDMA itself, only the particles with the dry
67 diameter under 350nm are usually observed. The $f(\text{RH})$ is calculated as the ratio of the
68 scattering coefficient at a certain RH to the corresponding dry scattering coefficient
69 (reference RH<40%), which can be measured by a dual nephelometer system (Covert
70 et al., 1972; Rood et al., 1985). In recent years, the κ -Köhler theory has been widely
71 used to describe the hygroscopic properties of aerosols (Petters and Kreidenweis, 2007),
72 in which all of the chemical composition dependent variables were merged into a
73 hygroscopicity parameter (κ). It facilitates the intercomparison of particle
74 hygroscopicity obtained by different equipment or different relative humidity. It is
75 applicable to both the single-component and the multicomponent aerosol particles.
76 Knowing the κ , dry diameter, and the relative humidity (RH), the $g(\text{RH})$ and liquid
77 water content (LWC) of aerosol particles could be calculated by using the κ -Köhler
78 equation.

79 The $f(\text{RH})$ synthetically reflects the impact of water uptake on aerosol particle size,
80 morphology, refractive index, etc., and consequently, on the scattering coefficient.
81 Firstly, the $f(\text{RH})$ results can be directly applied to the establishment of a



82 parameterization scheme for light extinction. The parameterization scheme could be
83 built based on chemical compositions (DeBell et al., 2006; Pitchord et al., 2007), mass
84 concentrations (Chen et al., 2015), or volume concentrations (Chen et al., 2012, 2014).
85 No matter which scheme is used, the effect of relative humidity on extinction or
86 scattering coefficient must be taken into account. Secondly, the $f(\text{RH})$ could be used to
87 calculate the hygroscopicity parameter (κ) combining with the measurement of particle
88 number concentration size distributions (PNSD) (Chen et al., 2014; Kuang et al., 2017).
89 This κ does not target specific particle size but represents the overall hygroscopicity of
90 the observed aerosol. Thirdly, the $f(\text{RH})$ results derived from a three-wavelength
91 nephelometer system have also been used for the calculation of aerosol liquid water
92 content (ALWC) (Kuang et al., 2018) and number concentrations of cloud condensation
93 nuclei (Tao et al., 2018).

94 The humidified nephelometer system was firstly used by Pilat and Charlson (1966)
95 to measure the effect of humidity on the light scattering and the size of NaCl particles.
96 Covert et al. (1972) then used a similar system to study the hygroscopic and/or
97 deliquescence effects which were dependent upon relative humidity for pure particles.
98 Rood et al. (1985) improved the humidified nephelometers, which could control the RH
99 more precisely. In the past two decades, Carrico et al. (2000), Day et al. (2000),
100 Koloutsou-Vakakis et al. (2001), Fierz-Schmidhauser et al. (2010), Liu et al. (2016),
101 etc. further improved the humidified nephelometers to accurately measure the $f(\text{RH})$
102 and study the light scattering enhancement characteristics of aerosol. While all the
103 humidified nephelometer system in the previous studies could only observe the $f(\text{RH})$
104 below 90% RH because there would be a $\pm 5\%$ error in the RH measurement in the
105 optical chamber of nephelometer when the RH was humidified over 90%. However, for
106 a hygroscopic particle, its particle size or scattering cross section will increase sharply
107 when the relative humidity exceeds 90%. Accordingly, the results above 90% RH would
108 be quite important for $f(\text{RH})$ curve fitting and light scattering calculating.

109 As we all know, the North China Plain (NCP) is the most severe area of air pollution
110 in China. The aerosol hygroscopicity is an important basis for the studies of
111 atmospheric visibility, radiative forcing, and aerosol secondary formation in this area.
112 The HTDMA had been used in some campaigns in NCP (Massling et al., 2009; Meier
113 et al., 2009; Liu et al., 2011; Wang et al., 2018). As mentioned above, the $g(\text{RH})$
114 observations using HTDMA had been basically limited to 350 nm, and could not be set
115 at multiple RH points. In comparison, the overall hygroscopicity of ambient aerosol at
116 several continuous RH points could be obtained by the $f(\text{RH})$ observation, which would
117 also be used to further calculate the hygroscopicity parameter and ALWC. In recent
118 years, the $f(\text{RH})$ observation has been carried out in only a few studies in NCP (Yan et
119 al., 2009; Pan et al., 2009; Chen et al., 2014; Kuang et al., 2017). In these studies, either
120 the temporal resolution of $f(\text{RH})$ was low, or the observation period was short. Moreover,
121 the $f(\text{RH})$ was limited to 90% RH in all these studies.

122 At present, some numerical pollution prediction systems have been established in the
123 region of NCP. However, the prediction of visibility is short of good means and methods.
124 Thus, an accurate, easy to use, and seasonal representative visibility parameterization
125 scheme is very critical and urgent. Furthermore, owing to the significant changes in the



126 pollution sources, the physical and chemical characteristics of aerosols in NCP have
127 also changed evidently in recent years. Above all, the research on aerosol
128 hygroscopicity is still not enough for NCP. Further study is required, especially in the
129 area of $f(\text{RH})$. In this study, an improved high-resolution humidified nephelometer
130 system was used to observe the $f(\text{RH})$ of $\text{PM}_{2.5}$ for a wide RH range between 30%-96%
131 in the urban area of Beijing over three seasons (winter, summer, and autumn). The main
132 objectives of this article are to characterize the variations of $f(\text{RH})$ and other optical
133 parameters under different seasons and different pollution levels, and set up the optimal
134 expressions of $f(\text{RH})$.

135 **2. Instruments and methods**

136 **2.1 Observation site**

137 The measurement campaign was performed at the Institute of Urban Meteorological
138 in the Haidian district (about 36m above the ground), which located in the northwest
139 urban area of Beijing, outside the third-ring road (39°56'N, 116°17'E). The sampling
140 site was located next to a high-density residential area, which has no significant
141 emissions from industrial in the surrounding neighborhood. Therefore, the observation
142 data could represent the air quality levels of the typical urban area of Beijing. The
143 observations were conducted in three seasons, 12nd Jan. to 14th Feb. for winter, 6th Jul.
144 to 21st Aug. for summer, and 30th Sep. to 13th Nov. for autumn.

145 **2.2 Instruments**

146 The $f(\text{RH})$ was measured by a dual-nephelometer system, one nephelometer for the
147 aerosol scattering coefficient under dry condition, and another nephelometer for the
148 humidified aerosol. The air flow first passed through a $\text{PM}_{2.5}$ inlet, and then was dried
149 by two tandem Nafion dryers (MD-700), which could reduce the RH of air flow lower
150 than 30%. The dried air was separated into two paths, one stream went directly into a
151 nephelometer, another stream was humidified by passing through a Gore-Tex tube,
152 which was set in a stainless steel tube. The interlayer between these two tubes is
153 circulating water headed by the water bath. The minutely scattering coefficients of dry
154 and humidified $\text{PM}_{2.5}$ under three wavelengths (450, 525, and 635nm) were
155 synchronously measured by these two nephelometers (Aurora 3000). In most studies,
156 there was only one water bath been used for humidifying. After a humidifying process,
157 it is necessary to wait for the water in the water bath to cool down enough for the next
158 process. Differently, in this study, two water baths were used. When one water bath was
159 heating up the water for humidifying, another water bath was cooling down the water
160 itself. After a humidifying process, the water bath with cool water would be switched
161 into the humidification pipeline. The use of two water baths could ensure that the
162 effective data of $f(\text{RH})$ is more than twice that of using only one water bath. The
163 temperature of water in the water bath was controlled by an automatic system to ensure
164 the humidifying effect.

165 Two combined RH and temperature sensors (Vaisala HMP110) were set at the inlet
166 and outlet of the wet nephelometer, respectively. The vapor pressures were calculated
167 by the sensor data, and the average value was considered as the vapor pressure in the
168 optical chamber. Thus, the humidified RH in the chamber could be calculated through
169 the derived vapor pressure and the temperature measured by the sensor in the chamber.



170 As mentioned above, the RH could not be accurately measured by a sensor when it is
171 above 90%. In order to accurately obtain higher relative humidity, the optical chamber
172 of wet nephelometer was deliberately cooled with the temperature lower than that at the
173 inlet and outlet. So the humidified RH could be higher than 95% in the chamber when
174 the RH at the inlet and outlet were lower than 90%. This method makes it possible to
175 observe the $f(\text{RH})$ under high RH. To avoid the vapor condensation and particle
176 activation, the upper limit of humidified RH in the optical chamber of wet nephelometer
177 was set to 97%, which made the effective data of $f(\text{RH})$ could reach RH of 96%. Each
178 humidifying process lasted about 50 minutes, and all the minutely average data were
179 automatically recorded by the control system. During the observation periods, these
180 two nephelometers were calibrated every ten days. Since two Vaisala sensors and two
181 nephelometers were all newly purchased and the relative error of vapor pressures was
182 always less than 1%, the sensors were not calibrated during the observation.

183 Other than the $f(\text{RH})$ measurement, the six-minute average $\text{PM}_{2.5}$ mass
184 concentrations were also measured by a continuous dichotomous ambient air monitor
185 (TEOM 1405DF). The sample filter and sample conditioner filter of this monitor were
186 replaced every 15 days or when the dust loading exceeded 70%. The five-minute
187 average absorption coefficient of $\text{PM}_{2.5}$ was monitored by a multiangle absorption
188 photometer (MAAP 5012). The quartz-fiber filter could be automatically changed when
189 the light transmission was less than 20%. In addition, the hourly water-soluble ions
190 (SO_4^{2-} , NO_3^- , Cl^- , NH_4^+ , Na^+ , K^+ , Mg^{2+} , and Ca^{2+}) of $\text{PM}_{2.5}$ and trace gases (HCl, HNO_3 ,
191 HNO_2 , SO_2 , and NH_3) were measured by an online analyzer (MARGA). In addition,
192 the aerosol number concentration distribution (SMPS3938+APS3321), and size-
193 resolved chemical compositions (MOUDI 122) were also synchronously measured
194 during these three observation periods.

195 In this paper, we mainly focus on the discussions of $f(\text{RH})$. In the near future, the
196 aerosol hygroscopicity would be comprehensively evaluated by making a use of all the
197 results from the above-mentioned observations and would be published in the following
198 papers.

199 2.3 Methods

200 The scattering Ångström exponent (Ångström, 1930) between 450 and 635nm (A_{450-}
201 $_{635}$) at dry condition ($\text{RH}<30\%$) characterizes the wavelength dependence of aerosol
202 scattering coefficients and was calculated using the scattering coefficient at
203 wavelengths of 635 and 450 nm by the following equation:

$$204 \quad A_{450-635} = (\log \sigma_{\text{sp}}(450) - \log \sigma_{\text{sp}}(635)) / (\log 635 - \log 450) \quad (1)$$

205 The $\text{PM}_{2.5}$ absorption coefficient was measured in the wavelength of 670nm by MAAP.
206 In order to facilitate comparison, we transform the absorption coefficient of 670nm into
207 that of 525nm according to the assumption that absorption is inversely proportional to
208 wavelength (Bond and Bergstrom, 2006; Liu et al., 2018). Thus, the SSA at 525nm
209 ($SSA_{525\text{nm}}$) was the proportion of the scattering coefficient to the sum of scattering
210 coefficient and absorption coefficient.

211 When the air is very clean, the relative change and fluctuation of the particle
212 concentration in the ambient air would be more intense. Furthermore, the airflow into
213 two nephelometers could not be completely synchronized during the observation. The



214 larger relative error of scattering coefficient in two nephelometers could lead to greater
215 fluctuations in $f(\text{RH})$ during the humidifying cycle when under clean conditions
216 because the $f(\text{RH})$ is the ratio of the scattering coefficients from humidified and dry
217 nephelometers. Thus, the $f(\text{RH})$ points with dry scattering coefficient at 525nm less than
218 50 were removed from the fitting of the $f(\text{RH})$ curves. In this paper, the $f(\text{RH})$
219 discussions are all based on the data of 525nm if not specifically pointed out.

220 The $\text{PM}_{2.5}$ concentrations are classified into three groups with $0\sim 35 \mu\text{g m}^{-3}$, $35\sim 75$
221 $\mu\text{g m}^{-3}$, and $>75 \mu\text{g m}^{-3}$, representing very clean, moderately clean, and polluted
222 conditions referring to the AQI grading standard of China, respectively. The reason for
223 this division is mainly based on the characteristics of $f(\text{RH})$ data.

224 3. Results and discussions

225 3.1 Overview of the optical properties and $f(\text{RH})$ of $\text{PM}_{2.5}$

226 Fig. 1 shows an overview of the hourly averaged light scattering coefficients
227 ($\sigma_{\text{sca},525\text{nm}}$), absorption coefficients ($\sigma_{\text{ap},630\text{nm}}$), single scattering albedo ($\text{SSA}_{630\text{nm}}$), and
228 scattering Ångström exponent ($\text{Å}_{450-635}$) as well as $f(\text{RH})$ at $\text{RH}=80\%$ ($f(80\%)$) for $\text{PM}_{2.5}$.
229 The average values of optical parameters and $f(80\%)$ in different seasons and under
230 different $\text{PM}_{2.5}$ pollution levels are listed in Table 1.

231 The $\text{PM}_{2.5}$ pollution was heaviest in the winter observation period and lightest in
232 summer. The scattering coefficient and absorption coefficient also show the same trends.
233 Single scattering albedo is one of the most important parameters in estimating of the
234 direct aerosol radiative forcing. The SSA_{525} increased with the aggravation of $\text{PM}_{2.5}$
235 pollution in all three seasons, indicating that the components with strong scattering
236 ability, such as secondary ions, increased significantly during the pollution process. The
237 wind rose of SSA_{525} in Figure 2 also indicates that the higher SSA_{525} values generally
238 occurred under the southerly wind condition which was often accompanied by higher
239 $\text{PM}_{2.5}$ concentrations.

240 Scattering Ångström exponent is generally regarded as an indicator of particle size.
241 In winter, lower $\text{Å}_{450-635}$ was observed in less polluted conditions. As depicted in Figure
242 2, the clean conditions in winter occurred mainly in the case of northwest wind with
243 relatively higher wind speed, which led to a greater proportion of larger particles such
244 as crustal dust in the air. Conversely, the $\text{Å}_{450-635}$ in polluted condition was lowest in
245 summer. Compared with winter, the relative humidity in summer was much higher,
246 especially under the condition of pollution, which could make particle collision and
247 coagulation easier to occur Guo et al (2014).

248 According to the ZSR (Zdanovskii-Stokes-Robinson) assumption (Zdanovskii, 1948;
249 Stokes and Robinson, 1966), the κ value of a multicomponent particle is equal to the
250 volume weighted average of each component. As depicted in Figure 1, the $f(80\%)$ at
251 525nm is highly correlated with the fractions of all the water-soluble ions to $\text{PM}_{2.5}$.
252 Owing to the proportion of hygroscopic components increased, $\text{PM}_{2.5}$ had higher $f(80\%)$
253 in the polluted conditions. The standard deviations (SD) of $f(80\%)$ also indicate that the
254 changes of fractions of hygroscopic components were relatively small when the $\text{PM}_{2.5}$
255 over $35\mu\text{g m}^{-3}$ in summer and autumn. The wind roses obviously reveal the differences
256 in the hygroscopicity and chemical compositions of $\text{PM}_{2.5}$ from different directions.
257 The $f(80\%)$ values in this study are in agreement with the range of values reported in



258 some other studies of the North China Plain (Table 2). Overall, the diurnal variation of
259 $f(80)$ is not obvious, and the average $f(80)$ at 12 to 16 pm was slightly higher (Figure
260 3).

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Table 1

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Table 2

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Figure 1

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Figure 2

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Figure 3

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272 3.2 Parameterization schemes of $f(\text{RH})$

273 To better describe the dependence of $f(\text{RH})$ on RH, many different empirical
274 expressions have been applied in previous studies to fit the $f(\text{RH})$ measurements.
275 Kotchenruther et al. (1999) proposed that different fitting equations should be used
276 according to the observed curve structure. For monotonic curves in which $f(\text{RH})$ varies
277 smoothly with RH, they proposed the use of equation reported by Kasten (1969) and its
278 variants. The most commonly used equation is the one-parameter fit equation (Hänel,
279 1980; Gassó et al., 2000; Brock et al. 2016) and (Kotchenruther and Hobbs, 1998;
280 Carrico et al., 2003; Zieger et al., 2011; Chen et al., 2014). For deliquescent curves,
281 Kotchenruther et al. (1999) introduced a more complex equation, and more detailed
282 information and fitting equations could be found in Titos et al. (2016).

283 In this work, four commonly used empirical parameterization schemes were chosen
284 to describe the monotonic curves of $f(\text{RH})$ variation:

$$285 \quad f(\text{RH})=a (1-\text{RH}/100)^{-\gamma(\text{RH}/100)} \quad \text{Chen et al. (2014)} \quad (2)$$

$$286 \quad f(\text{RH})=a (1-\text{RH}/100)^{-\gamma} \quad \text{Kasten (1969)} \quad (3)$$

$$287 \quad f(\text{RH})=1+a (\text{RH}/100)^{\gamma} \quad \text{Kotchenruther and Hobbs, (1998)} \quad (4)$$

$$288 \quad f(\text{RH})=1+a (\text{RH}/(100-\text{RH})) \quad \text{Brock et al. (2016)} \quad (5)$$

289 where γ parameterizes the magnitude of the scattering enhancement, which is not
290 affected by the RH. The comparison of the fitting results for different expressions is
291 shown in figure S1 to S3. According to the fitting results, R^2 , and the comparison
292 between fitting $f(80\%)$ values and measured ones, we finally choose the Eq. (2) to
293 describe the scattering enhancement due to monotonic hygroscopic growth.

294

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Figure 4

296

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Figure 5

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Figure 6

300

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Table 3



302

303 In our previous studies, the pollution of particulate matter was usually classified into
304 three categories through two threshold values of $PM_{2.5}$ concentrations, $75\mu\text{g}\cdot\text{m}^{-3}$ and
305 $150\mu\text{g}\cdot\text{m}^{-3}$, for clean, moderately polluted, and heavily polluted conditions. In this study,
306 we found that the scatter points of $f(\text{RH})$ were quite concentrated near the fitting curves
307 when the $PM_{2.5}$ was above $75\mu\text{g}\cdot\text{m}^{-3}$, and the fitting parameters were markedly different
308 for conditions of $PM_{2.5}$ above and under $35\mu\text{g}\cdot\text{m}^{-3}$. Consequently, we use $35\mu\text{g}\cdot\text{m}^{-3}$ and
309 $75\mu\text{g}\cdot\text{m}^{-3}$ as boundaries to classify the different pollution levels in this study. Fig. (4) to
310 Fig. (6) show the fitting $f(\text{RH})$ curves under very clean, moderately clean, and polluted
311 conditions for three seasons. Table 3 shows the fitting results from this work and a
312 previous study conducted in NCP, which used the same equation expression (Chen et
313 al., 2014). Except for the very clean condition in winter, the fitting values of a are all
314 near 1.0 and the fitting $f(\text{RH})$ curves are all above 1.0 for RH above 30%. However, the
315 fitting $f(\text{RH})$ curve shown in Fig 4a is apparently under the observed data points and
316 lower than 1.0 when RH under 40%. It means that the fitting curve of the very clean
317 condition in winter would underestimate the $f(\text{RH})$ value when RH is below 40%. It
318 can be seen from Figure 5 and Figure 6 that the fitting R^2 increases and the points
319 become more concentrated along with the aggravation of $PM_{2.5}$ pollution in summer
320 and autumn. This is in accordance with the characteristics of the SD of $f(80)$ in Figure
321 3. Differently, the $f(\text{RH})$ points are still dispersed even under the polluted condition of
322 winter. This indicates that the chemical compositions of $PM_{2.5}$ when under the condition
323 of pollution were more stable in summer and autumn.

324 By comparing the fitting curves of different seasons and different pollution
325 conditions, it is found that the $f(\text{RH})$ curve of the very clean condition is lower than that
326 of other two pollution levels for each season. Except for the very clean condition in
327 winter and autumn, the fitting curves for the other seven conditions are relatively close,
328 especially when RH less than 90%. In the study of Chen et al. (2014), a segment fitting
329 with the critical RH of 60% is applied in the parameterization of $f(\text{RH})$. Comparing the
330 fitting results in Table 3, we find that the $f(\text{RH})$ values with RH above 80% from fitting
331 curves for clean and polluted conditions in Chen's study are evidently lower than those
332 from the respective curves for very clean and polluted conditions in this work. It
333 indicates that the scattering enhancement due to moisture uptake or hygroscopicity of
334 aerosol under high RH is probably higher than before.

335 In addition, the averaged $f(\text{RH})$ at 450nm, 525nm, and 635nm was also fitted
336 separately for each season. It is clear that the $f(\text{RH})$ showed a dependence on
337 wavelength, especially in summer. The averaged $f(\text{RH})$ increased with increasing
338 wavelength. Similar results were also obtained by Zhang et al. (2015) at Lin'an, China
339 and Zieger et al. (2014) at a regional continental research site in Melpitz, Germany.
340 However, we found that when under very clean conditions in winter and autumn, the
341 mean value of $f(80)$ at 450nm was higher than that at 525nm and 635nm (Fig. S4), and
342 it is more obvious in winter. Through further curve fitting of $f(\text{RH})$ at different
343 wavelengths for three seasons, it is also found that the $f(\text{RH})$ curve of 450nm was
344 evidently higher than that of 525nm and 635nm only under very clean conditions in
345 winter. Our previous work showed that the sulfates, nitrates, and ammonium (SNA)



346 were abundant in aerodynamic diameter of 0.18~1.0 μm on clean days in winter with
347 the mass median diameters (MMD) of SNA at about 0.45 μm . However, the SNA was
348 mainly concentrated in 0.32~1.8 μm with evidently higher MMDs on polluted days or
349 in other seasons (Zhao et al., 2017; Su et al., 2018). The high fraction of SNA in
350 particles below 500 nm might be responsible for the higher $f(\text{RH})$ at 450nm when under
351 very clean conditions in winter.

352 3.3 Uncertainty analysis for $f(\text{RH})$ measurements

353 As mentioned above, the humidified RH in the wet nephelometer was calculated
354 through the derived average vapor pressure and the temperature measured by the sensor
355 in the chamber. According to the differences in vapor pressure values from the sensors
356 at the inlet and outlet, a relative error of 0.5% could be calculated for the vapor pressure
357 data. And the mean absolute error of the temperature measurement in the nephelometer
358 was 0.2°C. Then, a Monte Carlo simulation was utilized to estimate the uncertainty of
359 calculated humidified RH values. New values of vapor pressure and temperature were
360 simulated by adding the uncertainties of 0.5% and 0.2°C to the observation data
361 (following a normal distribution). And new humidified RH could be calculated using
362 the simulated data, and this procedure was repeated 1000 times for each humidified RH
363 value. Then the average standard error of humidified RH in the wet nephelometer could
364 be calculated to be 0.85%.

365 Next, the Monte Carlo simulation was used again. The RH was assumed to range
366 from 20% to 90% with steps of 1, and assuming that γ ranges from 0 to 1 with steps of
367 0.01. The chosen interval for γ covers the particle types from non-hygroscopic aerosol
368 particles to very hygroscopic particles. Therefore, more than 7000 conditions were
369 simulated, and each condition corresponded to one set of RH and γ . For each condition,
370 the dry scattering coefficients with a wide range of 1 to 1000 Mm^{-1} were selected 5000
371 times as random numbers to present different atmospheric situations and aerosol loads.
372 The wet scattering coefficients were calculated associated with the previously selected
373 dry scattering coefficient and the $f(\text{RH})$ calculated by Eq. (2). A random error
374 (following a normal distribution with a standard variation of 0.85%) was also added to
375 the RH and the parameter a was set 1.0 when calculating the $f(\text{RH})$. According to the
376 manual of nephelometer Aurora 3000, the standard error of aerosol scattering
377 coefficient is 2.5%. Then we simulated the dry and wet scattering coefficients by
378 assuming that they both had an uncertainty of 2.5% (following a normal distribution).
379 Thus, the simulated $f(\text{RH})$ can be calculated again.

380 The mean and relative standard deviation of simulated $f(\text{RH})$ were calculated for each
381 RH and γ and are shown in Fig. 7. For aerosols in this work ($\gamma \sim 0.35$), $f(\text{RH})$ errors
382 were below 4% with RH lower than 85% while reached 9.7% when RH= 96%, which
383 can be regarded as a conservative estimation. And other unpredictable factors
384 contributing to $f(\text{RH})$ uncertainty have not been considered in this approach.

385 Figure 7

386 4. Conclusions

387 A wide-range (30%-96%) and high-resolution humidified nephelometer system was
388 developed and a measurement campaign was conducted to study the $f(\text{RH})$ for three
389



390 seasons in 2017. The $f(\text{RH})$ at higher RH had firstly been monitored and reported.

391 Based on the overview of the optical properties and $f(\text{RH})$ of $\text{PM}_{2.5}$, we found that
392 the higher SSA_{525} values generally occurred under the southerly wind components with
393 higher $\text{PM}_{2.5}$ concentrations. The $f(80\%)$ at 525nm of $\text{PM}_{2.5}$ was evidently higher under
394 the polluted conditions and highly correlated with the fractions of all the water-soluble
395 ions. The average $f(80)$ at 12 to 16 pm was slightly higher than that of other periods
396 from the average diurnal variations.

397 By comparing the fitting results and curves of four different empirical
398 parameterization schemes, it was found that one of the two-parameter fit equations can
399 better fit the observed $f(\text{RH})$ data. The fitting curves could be widely applied to the
400 studies of atmospheric visibility, radiative force, or liquid water content due to aerosol
401 moisture absorption. For summer and autumn, the $f(\text{RH})$ points of polluted conditions
402 were more concentrated near the fitting curves in the scatter plots with higher fitting R^2 .
403 And the fitting curve under the very clean condition was lower than that of other
404 conditions for each season. Compared with the fitting curves in the previous study, the
405 hygroscopicity of aerosol under higher RH has probably been enhanced. In summer,
406 the fitting $f(\text{RH})$ showed a significant dependence on wavelength and increased with
407 increasing wavelength for each pollution condition. Nevertheless, the $f(\text{RH})$ curve of
408 450nm was evidently higher than that of 525nm and 635nm under the very clean
409 condition in winter, due to the higher fraction of SNA in particles below 500 nm.

410 The uncertainties of $f(\text{RH})$ were simulated by considering all the predictable
411 uncertainties or errors during the measurement, which was below 10% for RH up to
412 96%.

413

414 *Data availability.* All data in this work are available by contacting the corresponding
415 author P. S. Zhao (pszhao@ium.cn).

416

417 *Author contributions.* P Z designed and performed this study. P Z and J D analyzed the
418 data and discussed the results. P Z prepared the manuscript and J D prepared all the
419 figures. X D and J S calibrated the device and collected the data.

420

421 *Competing interests.* The authors declare that they have no conflict of interest.

422

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- 606



607 **Table captions:**

608 Table 1. Average $\sigma_{\text{sca},525\text{nm}}(\text{dry})$, $\sigma_{\text{abs},525\text{nm}}(\text{dry})$, $\text{SSA}_{525}(\text{dry})$, $\text{Å}_{450-635}(\text{dry})$, and $f(80\%)$
609 under different pollution levels in three seasons.

610 Table 2. Comparisons of average $f(80\%)$ in different campaigns of NCP area.

611 Table 3. Comparisons of fitting parameters with another study using the same scheme.

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Table 1

		Entire observation periods	PM _{2.5} pollution levels (µg m ⁻³)		
			Very clean (PM _{2.5} ≤35)	Moderately clean (35 < PM _{2.5} ≤ 75)	Polluted (PM _{2.5} > 75)
Winter	σ _{sca,525nm} (dry)	287±359	35±19	205±69	504±158
	σ _{abs,525nm} (dry)	46±37	10±7	41±13	75±15
	SSA ₅₂₅ (dry)	0.82±0.06	0.79±0.04	0.83±0.03	0.87±0.02
	Å ₄₅₀₋₆₃₅ (dry)	1.27±0.20	1.18±0.13	1.42±0.15	1.45±0.15
	f(80%)	1.47±0.16	1.31±0.12	1.51±0.09	1.60±0.14
Summer	σ _{sca,525nm} (dry)	170±125	85±46	226±87	410±121
	σ _{abs,525nm} (dry)	27±13	20±10	30±11	40±12
	SSA ₅₂₅ (dry)	0.83±0.09	0.79±0.09	0.87±0.05	0.91±0.02
	Å ₄₅₀₋₆₃₅ (dry)	1.34±0.23	1.42±0.22	1.30±0.22	1.18±0.21
	f(80%)	1.54±0.16	1.50±0.19	1.60±0.08	1.62±0.06
Autumn	σ _{sca,525nm} (dry)	261±243	57±54	241±62	564±169
	σ _{abs,525nm} (dry)	38±24	17±12	40±16	61±18
	SSA ₅₂₅ (dry)	0.83±0.09	0.75±0.09	0.85±0.05	0.90±0.03
	Å ₄₅₀₋₆₃₅ (dry)	1.17±0.27	1.16±0.27	1.30±0.20	1.06±0.27
	f(80%)	1.53±0.11	1.44±0.14	1.58±0.08	1.57±0.04

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Table 2

Study area (Campaign)	Periods	Aerosol pollution levels	f(RH=80%)	Wavelength (nm)	Reference
The rural site of Beijing	24 April–15 May 2006	Clean	1.31 ± 0.03	525	Pan et al. (2009)
		Urban pollution	1.57 ± 0.02		
		Special case	2.21		
SDZ, Beijing, a rural site	December	Relatively clean	1.16	525	Yan et al. (2009)
		Relatively polluted	1.34		
CAMS, Beijing, an urban site	2005	Relatively clean	1.2	525	(2009)
		Relatively polluted	1.48		
Wuqing, Tianjin	October 2009 to late January 2010	Clean	1.46 ± 0.15	550	Chen et al. (2014)
		Polluted	1.58 ± 0.19		
Wangdu, suburban district of North China Plain	4 June 2014 - 14 July 2014	Entire campaign deliquescent phenomena	1.8 (1.1-2.3) 2.0 (1.7-2.3)	550	Kuang at al. (2017)

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Table 3

	Periods	a	γ	Reference
Very clean		0.930	0.329	
Moderately clean	12 Jan.-14 Feb., 2017	0.971	0.372	
Polluted		0.988	0.356	
Very clean		0.972	0.355	
Moderately clean	6 July-21 Aug., 2017	0.980	0.362	This work
Polluted		0.984	0.371	
Very clean		0.979	0.334	
Moderately clean	30 Sep. to 13 Nov., 2017	1.002	0.344	
Polluted		1.014	0.332	
Entire campaign	RH<60%	1.02	0.21	
	RH≥60%	1.08	0.26	
Clean	Oct.2009- Jan.2010	RH<60%	1.00	Chen et al. (2014)
		RH≥60%	1.00	
Polluted		RH<60%	1.03	
		RH≥60%	1.14	0.25

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629 **Figure captions:**

630 Figure 1. Time series of $\sigma_{\text{sca},525\text{nm}}(\text{dry})$, $\sigma_{\text{abs},525\text{nm}}(\text{dry})$, $\text{SSA}_{525}(\text{dry})$, $\text{Å}_{450-635}(\text{dry})$,
631 $f(80\%)$, and mass concentrations of water-soluble ions and their mass fractions over
632 all the sampling periods in three seasons.

633 Figure 2. Wind dependence of $\sigma_{\text{sca},525\text{nm}}(\text{dry})$, $\sigma_{\text{abs},525\text{nm}}(\text{dry})$, $\text{SSA}_{525}(\text{dry})$, $\text{Å}_{450-635}(\text{dry})$,
634 and $f(80\%)$ over three seasons; the shaded contour indicates the average of variables
635 for varying wind speeds and wind directions.

636 Figure 3. Diurnal variations of $f(80\%)$ in different seasons.

637 Figure 4. Fitting $f(\text{RH})$ curves under different pollution levels in winter.

638 Figure 5. Fitting $f(\text{RH})$ curves under different pollution levels in summer.

639 Figure 6. Fitting $f(\text{RH})$ curves under different pollution levels in autumn.

640 Figure 7. Simulated $f(\text{RH})$ and its error (color scale) as a function of RH and the
641 hygroscopic parameter γ .

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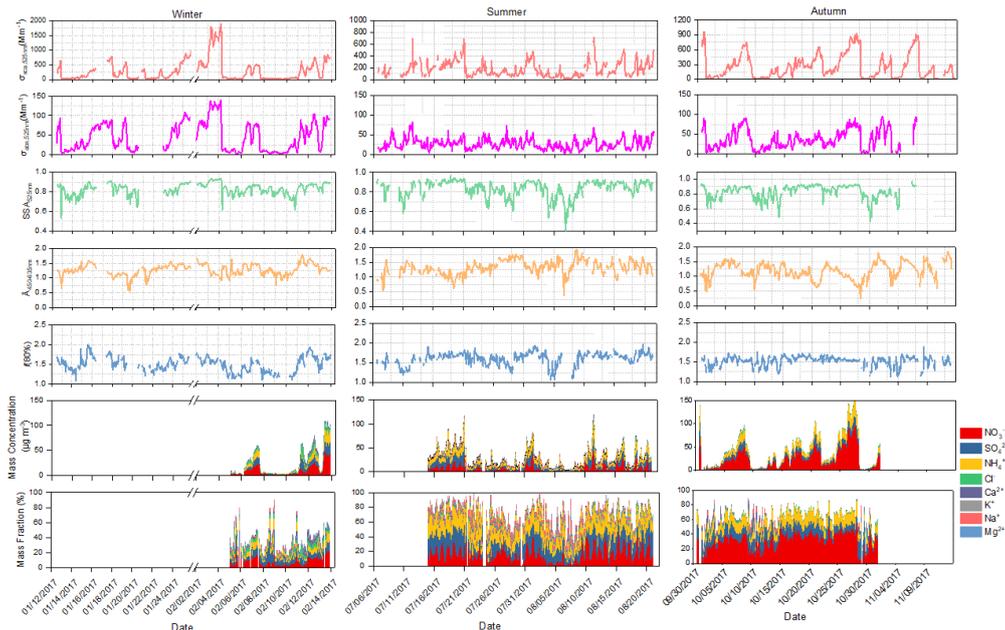
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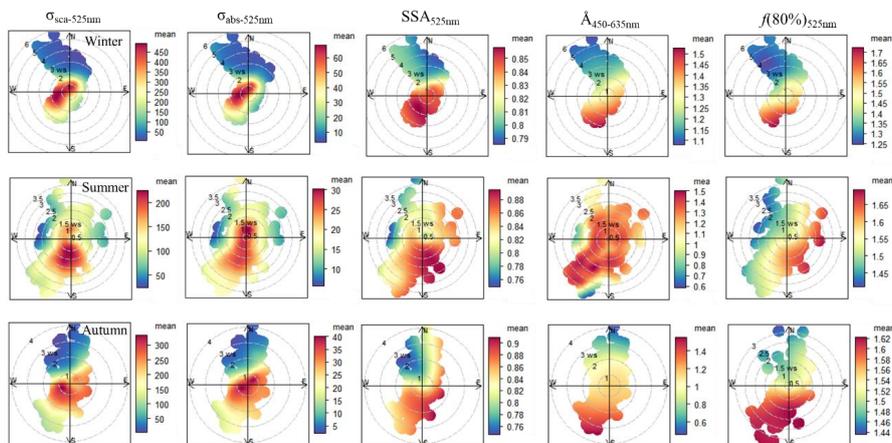
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Figure 1

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Figure 2

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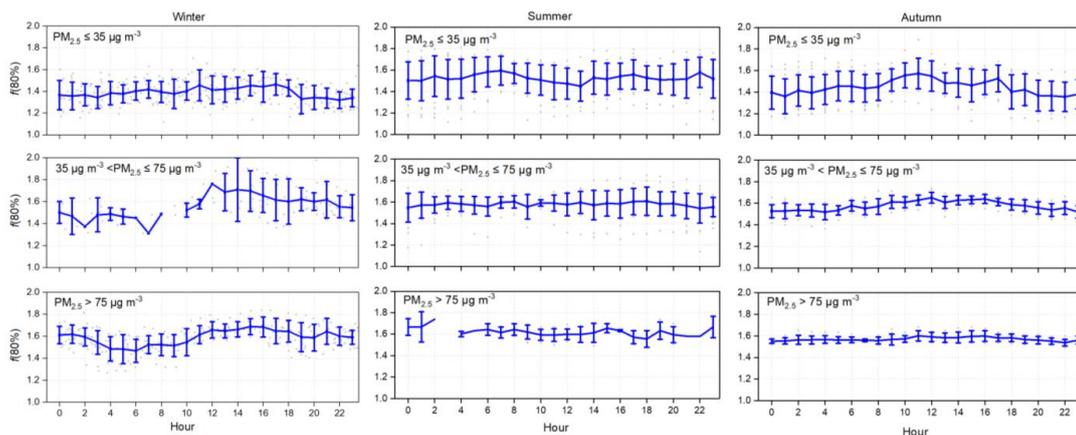
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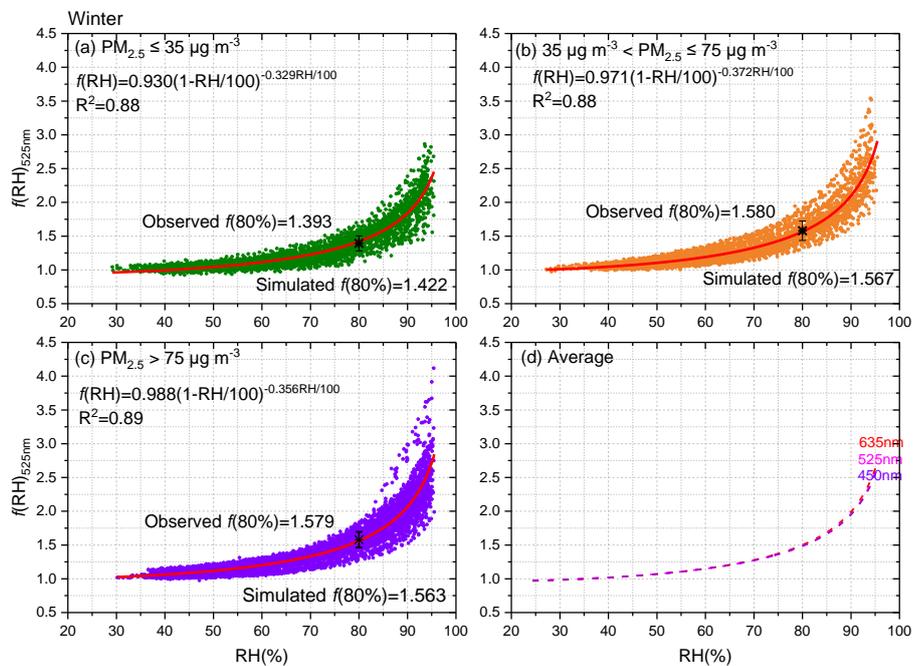


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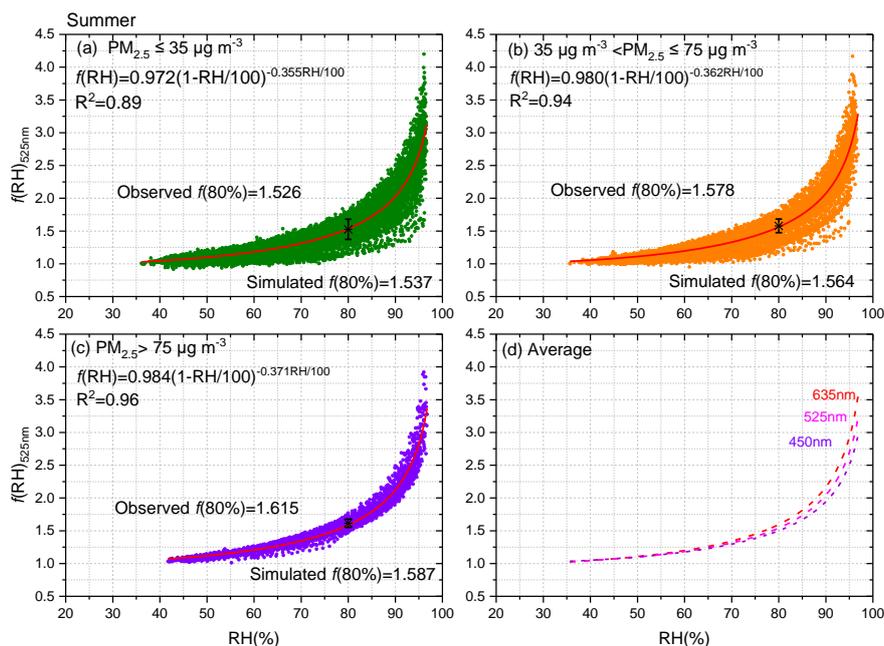
Figure 3



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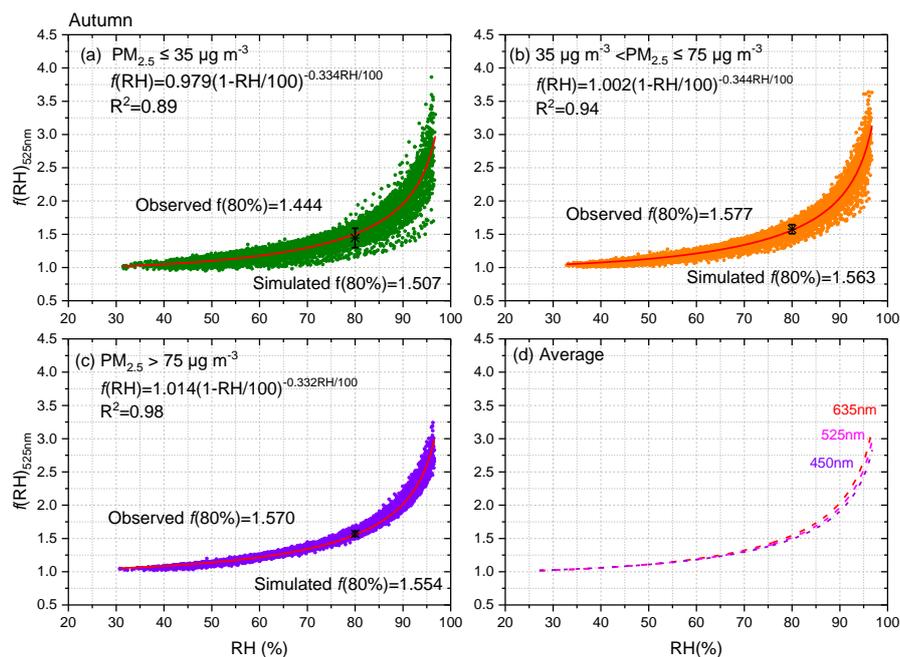
Figure 4



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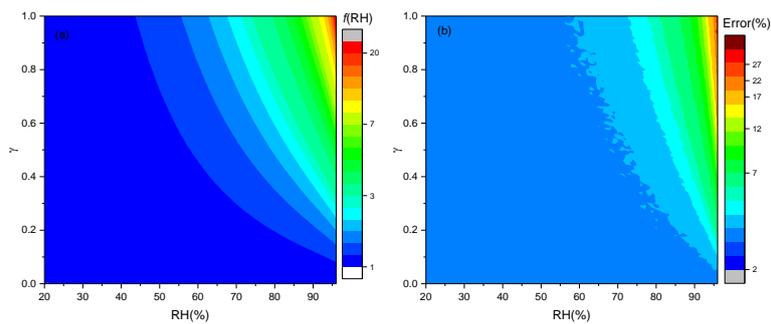
Figure 5



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Figure 6



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Figure 7