



1 **Formation and characteristics of secondary aerosols in an**
2 **industrialized environment during cold seasons**

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18 **Abstract.** Secondary aerosols including inorganic and organic components often
19 dominate the fine aerosol mass, it is thus important to elucidate the formation and
20 characteristics of these species. In this work, we measured the submicron aerosols
21 (PM₁) by using an Aerodyne high resolution soot-particle aerosol mass spectrometer
22 in suburban Nanjing, China. The site was surrounded by industry plants, and the
23 measurement was conducted during cold seasons (February-March 2015). We found
24 that under such environment, the PM₁ was predominantly comprised of secondary
25 species (on average 63.2% from ammonium sulfate and nitrate). Results show that
26 moisture plays a key role to enhance both nitrate and sulfate formations. The moisture
27 promotes the gas-particle partitioning and nocturnal heterogeneous production of
28 nitrate, while transformation of SO₂ into sulfate directly in aqueous phase is more
29 significant. The organic aerosol (OA) occupied ~1/4 of total PM₁ mass, and the



30 primary OA (POA) and secondary OA (SOA) contributions were almost equal. A
31 specific industry-related OA was separated and a modified graphical method was
32 introduced to describe the evolution of OA. Results further show that the most
33 abundant OA factor, which is the one with highest oxidation degree, is also mainly
34 driven by aqueous-phase processing, while the other two less oxygenated SOA factors
35 are mainly governed by photochemical processing. Peak sizes of sulfate, nitrate and
36 OA all shifted towards larger sizes with the increases of relative humidity, reflecting
37 the effects of aqueous-phase processing too. Aqueous-phase driven secondary
38 aerosols were found to be very important in enhancing the PM₁ pollution, while
39 photochemical processed SOA was important to OA pollution, leading to a fresher OA
40 at higher OA concentrations. We further demonstrated influences of the
41 aqueous-phase processing and photochemical processing on formation of secondary
42 aerosols by using two typical cases, respectively. This paper highlights the importance
43 of aqueous-phase chemistry on sulfate and nitrate formations, and that different
44 portions of SOA can be dominated by different mechanisms in an industrialized
45 environment.

46

47 1. Introduction

48 Aerosol particles can significantly affect the earth's climate (e.g., Pöschl,
49 2005; Carslaw et al., 2010), air quality (e.g., Chan and Yao, 2008; Shen et al.,
50 2017; Heal et al., 2012) and human health (e.g., Shiraiwa et al., 2017; Pope and
51 Dockery, 2006; Hu et al., 2017), etc. However, such effects are highly uncertain, in a
52 large part due to varying physicochemical properties of the particles, including sizes,
53 compositions, and sources, etc. The secondary species, which often dominate the fine
54 particle mass (e.g., Zhang et al., 2007; Huang et al., 2014), are particularly less
55 understood. Formation of secondary components, including both inorganic and
56 organic species, is very complex and dependent heavily upon atmospheric
57 environments and meteorological conditions. For example, recent studies demonstrate
58 that nitrate formation can be governed by multiple mechanisms in different seasons
59 and locations of China (Ge et al., 2017a; Yang et al., 2017). For sulfate, it is well



60 known that the transformation of SO₂ into sulfate can occur in both gas-phase and
61 aqueous-phase via reactions with multiple oxidants (Seinfeld and Pandis, 2016).
62 Similarly, secondary organic aerosol (SOA) can be produced by both gas-phase and
63 aqueous-phase (or multiphase/heterogeneous) reactions too (e.g., McNeill,
64 2015;Herrmann et al., 2015;Blando and Turpin, 2000;Lim et al., 2010). For examples,
65 Zhang et al. (2018a) points out that both photochemical and aqueous-phase reactions
66 can be important for the secondary organic carbon (SOC) in Shanghai during
67 summertime, and aqueous-phase processing is more important to the SOC nighttime
68 formation; Saffari et al. (2016) also shows the nighttime SOC formation in Los
69 Angeles, while Ye et al. (2017b) shows that photochemical oxidation is important to
70 the daytime SOC formation in Changzhou, China.

71 Due to a large variety of possible gaseous organic precursors, the formation
72 pathways, yields and properties of SOA are complicated and much less clear
73 compared with those of secondary inorganic aerosol (SIA). A large number of
74 laboratory studies have been conducted to elucidate the characteristics of SOA
75 produced by gas-phase oxidation (gasSOA) from certain precursors and the relevant
76 mechanisms are continuously implemented into models (Ervens, 2015;Hallquist et al.,
77 2009). Nevertheless, current models still cannot accurately reproduce the SOA mass
78 or other properties, and the SOA produced in aqueous-phase (cloud/fog drops or
79 aqueous aerosols) (aqSOA) is postulated to be a possible missing portion that can help
80 reconcile the discrepancies (Ervens, 2015;Ervens et al., 2011;Heald et al., 2005).
81 Correspondingly, studies regarding aqSOA formation have also made significant
82 achievements recently (Herrmann et al., 2015;George et al., 2015a). For example,
83 some volatile organic compounds, like isoprene (Liu et al., 2012), glyoxal and
84 methylglyoxal (Lim et al., 2013), pyruvic acid (Carlton et al., 2006), amines (e.g., Ge
85 et al., 2011b, a;De Haan et al., 2011), phenolic species (Yu et al., 2016;George et al.,
86 2015b;Yu et al., 2014;Sun et al., 2010), etc., are found to be effective aqSOA
87 precursors, and the aqSOA is significantly different from gasSOA in terms of product
88 identities, volatility, hygroscopicity and optical properties (e.g., Lim et al.,



89 2013;Ervens et al., 2011).

90 As ambient conditions are more variable than the lab-controlled conditions,
91 elucidation of secondary aerosol formation (including both SIA and SOA) and its
92 influencing factors is more sophisticated. Traditional filter-based studies are limited in
93 capturing the rapid evolution process, thus highly-time resolved measurements are
94 necessary for a better interpretation of the secondary aerosol formation in ambient air
95 (Wexler and Johnston, 2008). Recently, the Aerodyne Aerosol mass spectrometry
96 (AMS) technique (Canagaratna et al., 2007) has emerged as a powerful tool, as it can
97 quickly determine the concentrations, sizes and chemical compositions of fine
98 aerosols (typically for submicron aerosols, PM₁). In combination with factor analysis
99 (e.g., Zhang et al., 2011), it is able to resolve a few types of SOA, which may reflect
100 different SOA formation mechanisms and/or evolution processes. For example, the
101 effects of aqueous-phase and photochemical processing on the SOA formation in
102 Beijing were discussed through analyses of multiple datasets (Xu et al., 2017a; Sun et
103 al., 2016; Sun et al., 2013a). Formation of the aqSOA from AMS was also observed in
104 other locations (e.g., Ge et al., 2012; Gilardoni et al., 2016; Li et al., 2013). Overall,
105 these AMS results have greatly advanced our understanding on SOA (e.g., Jimenez et
106 al., 2009; Li et al., 2017; Spracklen et al., 2011; Philip et al., 2014).

107 As is well known, Eastern China is frequently experiencing severe haze
108 pollutions, and Yangtze River Delta (YRD) region is one of the heavily polluted
109 regions (e.g., Hu et al., 2014). Previously, the AMS has been applied for online
110 characterization of the fine aerosols in Nanjing (Shen et al., 2014; Zhang et al.,
111 2015a; Zhang et al., 2016; Zhang et al., 2017; Wang et al., 2016b) and Hangzhou (Li et
112 al., 2018) of this region. The offline AMS technique is also developed to analyze filter
113 samples collected in Changzhou (Ye et al., 2017c; Ye et al., 2017a), and Yangzhou (Ge
114 et al., 2017b) in the YRD region. In this study, we conducted a field campaign by
115 using the most advanced AMS version, soot-particle aerosol mass spectrometer
116 (SP-AMS) (Onasch et al., 2012) in suburban Nanjing during February - March 2015.
117 The sampling site was near an industry zone, and indeed we reported earlier the



118 occurrence of fullerene soot, that was linked with industry emissions (Wang et al.,
119 2016c). In this work, we focused on analyzing the characteristics of secondary
120 aerosols in this special industrialized environment.

121

122 **2. Experimental Methods**

123 **2.1 Sampling site and instrumentation**

124 The sampling site and instrumentation was described in Wang et al. (2016c), thus
125 only a brief summary is provided here. The site resided in the Campus of Nanjing
126 University of Information Science and Technology (32°12'20.82"N, 118°42'25.46"E),
127 and the measurement period was February 20 to March 23, 2015. Particularly, the site
128 was located west/southwest of an industrial zone (mainly petrochemical, chemical,
129 iron and steelmaking, and power plants), close to the residential area and a few
130 arterial roads (within a radius of a few kilometers) (Fig. 1). Therefore, the site was
131 probably influenced by the mixed emissions from industry, traffic and cooking, etc.

132 Our main instrument was SP-AMS, which was carefully tuned and strictly
133 calibrated for both mass quantification and sizes following the standard protocols
134 (details in Wang et al. (2016c)). The SP-AMS was operated with a dual-vaporizer
135 setup (with both laser and tungsten vaporizers), and the laser was switched alternately,
136 thus we were able to determine the chemical compositions and size distributions of
137 both non-refractory species (ammonium, sulfate, nitrate, chloride and organics) and
138 refractory black carbon (*r*BC). Concentrations of PM_{2.5} and gaseous species (CO,
139 NO₂, SO₂ and O₃) were acquired from the nearest environmental monitoring site. The
140 meteorological parameters, including air temperature (T), relative humidity (RH),
141 wind speed (WS), wind direction (WD), solar radiation and visibility were provided
142 by a meteorological station ~50m away from our site.

143

144 **2.2. Data analyses**

145 The SP-AMS data were processed by using the Igor Pro-based (Wavemetrics)
146 standard ToF-AMS analysis toolkit SQUIRREL version 1.59D and PIKA version



147 1.19D (Sueper, 2015). All mass concentrations were calculated from the high
148 resolution fitting of V-mode data (a mode sensitive to mass changes). The *r*BC
149 concentrations were from V-mode with both laser and tungsten vaporizers on, while
150 concentrations of other non-refractory species were from V-mode with the tungsten
151 vaporizer only. The calibrated relative ionization efficiencies (RIEs) were used to
152 account for the different instrument responses to different species. Also, the
153 composition-dependent collection efficiencies (CE) (Middlebrook et al., 2012) were
154 applied to consider the particle loss due to incomplete transmission, and particle
155 bouncing from the tungsten heater. All these data were averaged into hourly data
156 when comparing with the meteorological parameters or gaseous species. The data
157 reported are at local time (Beijing time).

158 The high chemical resolution of SP-AMS allows us to separate different ions and
159 derive the elemental ratios of OA including oxygen-to-carbon (O/C),
160 hydrogen-to-carbon (H/C), nitrogen-to-carbon (N/C) ratios, and organic mass to
161 organic carbon (OM/OC) ratios. The method proposed by Canagaratna et al.
162 (2015)(referred to as I-A method) was used here unless otherwise stated. The I-A
163 method is an update of the Aiken-ambient (A-A) method (Aiken et al., 2008), which
164 improved the calculation of O/C and H/C ratios. Results from I-A method correlated
165 very well with those from A-A method (Fig. S1 in the supplement), but increased the
166 O/C, H/C and OM/OC on average by 27%, 9% and 8.5%, respectively.

167 Positive matrix factorization (PMF) (Paatero and Tapper, 1994) and the PMF
168 Evaluation Toolkit version 2.08D (Ulbrich et al., 2009) were applied on the high
169 resolution mass spectra of OA acquired under W-mode (a mode with high mass
170 resolution, ~4000 in this work) with the dual-vaporizer setting. Six OA factors were
171 identified, including three primary factors relevant with traffic (HOA), cooking
172 (HOA), industry (IOA), and three secondary OA factors which are a semi-volatile
173 oxygenated OA (SVOOA), a low-volatility OA (LVOOA) and a local secondary OA
174 (LSOA). Procedures of the PMF analyses, justifications and diagnostics of the
175 optimal PMF solution, and the high resolution mass spectra (HRMS) of the factors



176 were presented in details in Wang et al. (2016c). In this work, we focused on
177 analyzing the features and behaviors of these OA factors, in particular the unique
178 primary IOA and three SOA factors.

179

180 **3. Results and discussion**

181 **3.1. Overview of the PM₁ characteristics**

182 Figure 2 shows the time series of meteorological parameters (T, RH, solar
183 radiation, WS, WD), concentrations of gaseous species (NO₂, O₃, SO₂ and CO),
184 fractional contributions of different PM₁ components (sulfate, nitrate, chloride,
185 ammonium, organics and *r*BC) to the total PM₁, and different OA factors (HOA, COA,
186 IOA, LSOA, SVOOA and LVOOA) to the total OA. The weather was overall humid
187 during the sampling period with an average RH of ~70%. There were also a few
188 precipitation events, for example, intermittently on February 28-March 1 and March
189 17-19, unfortunately the precipitation data were not recorded due to malfunction of
190 the meteorological meter. The wind was not very strong (average 1.4 m/s) and mostly
191 blew from east/northeast directions (Fig. S2). As a number of industrial plants located
192 east/northeast of our site, the measured PM was expected to be influenced by
193 industrial emissions. The campaign acrossed the later winter and early spring, thus the
194 temperature varied significantly from 0°C to 21°C with a mean of 8.5°C. The later
195 days (March 13-23) were much warmer than the earlier days (12.5°C vs. 6.0°C).

196 The PM₁ concentrations varied dynamically from 8.4 to 180.5 µg/m³, with an
197 average of 46.3 µg/m³, and was dominated by inorganic components (68.4%) (Fig.
198 3a). Such mass contribution from inorganic species in PM₁ was higher than most
199 AMS measurements in both urban and rural sites of China (Li et al., 2017) and other
200 countries (Jimenez et al., 2009), likely reflecting the SO₂ and NO_x emissions
201 enhanced by industry at this specific location. *r*BC contributed 6.1% of PM₁ mass.
202 Organics, on average, was still the most abundant component but only took up 25.5%
203 of PM₁ mass. The average contribution from POA (=HOA+COA+IOA) (52%) was
204 almost equal to that of SOA (=LSOA+SVOOA+LVOOA) (48%) (Fig. 3b). Taking



205 together, the secondary components (SIA and SOA) occupied ~82% of the PM₁ mass.
206 This fraction is higher than those observed in PM_{2.5} in a few Chinese megacities
207 during heavy haze periods (30-77%) (Huang et al., 2014). It points out that even in an
208 environment that is expected to have significant primary emissions, secondary species
209 still plays a major role to the aerosol pollution, indicating that the precursors (SO₂,
210 NO₂ and SOA precursors) can be efficiently oxidized in such atmosphere.

211 The average size distributions of different species are shown in Fig. 3c. The
212 major inorganic species all peaked in the accumulation mode (~650 nm D_{va}),
213 representing their behaviors as secondary species. Relatively, the organics had a wider
214 size distribution and peaked at a smaller size (~550 nm D_{va}), indicating it was a
215 mixture of both primary and secondary species. Size distribution of *r*BC was
216 significantly different from other components (peak D_{va} of 250-500 nm), as it was
217 predominantly originated from primary sources.

218 We also compared the mass concentrations of PM₁ determined by SP-AMS with
219 the PM_{2.5} concentrations from the nearest monitoring station in Fig. 3d. Generally,
220 they correlated well with each other ($r^2=0.70$), and the PM₁ on average occupied ~83%
221 of PM_{2.5}. This PM₁/PM_{2.5} fraction is higher than those previously observed in urban
222 Nanjing, for example, 54% during springtime (Wang et al., 2016b) and 63% during
223 wintertime (Zhang et al., 2016), showing a more significant contribution from small
224 particles in an environment affected by industry than the average urban case.

225

226 3.2 Sulfate and nitrate formations

227 As shown in Fig. 3a, sulfate and nitrate were abundant in PM₁. By assuming they
228 were associated with ammonium, (NH₄)₂SO₄ and NH₄NO₃ together would dominate
229 the PM₁ mass (63.2%). In Fig. 4, we presented the variations of sulfate and nitrate
230 concentrations as functions of RH and the odd oxygen ($O_x=O_3+NO_2$) over the entire
231 sampling period. Typically, RH is an indicator of atmospheric moisture and is relevant
232 with aqueous-phase/heterogeneous reactions (Tie et al., 2017), while O_x can be used
233 to represent photochemical activities (Herndon et al., 2008). From Fig. 4a, both



234 sulfate and nitrate concentrations increased substantially with the increase of RH,
235 especially at $RH > 65\%$. Note the decreases at $RH > 90\%$ were probably due to the
236 scavenging effects of precipitation (we have no exact precipitation data so cannot
237 accurately eliminate such influences). On the other hand, sulfate and nitrate both
238 presented no clear increases with the increase of O_x (Fig. 4b). The image plots which
239 describe the dependences of nitrate (Fig. 4c) and sulfate (Fig. 4d) on RH and O_x , also
240 showed that high mass concentrations of sulfate and nitrate appeared mainly in the
241 regimes with $RH > 65\%$, but distributed evenly across the changes of O_x . These results
242 highlight a more significant role of moisture in enhancing the formation of both
243 sulfate and nitrate than that of photochemical processing. Similar effects of RH on
244 sulfate and nitrate were also observed in urban Jinan, China (Wang et al., 2012).

245 As is well known, ammonium nitrate is semi-volatile and water can enhance its
246 thermodynamic gas-particle partitioning and dissolution in the particle phase. Our
247 results clearly reveal that this “thermodynamically-driven” mechanism (Ge et al.,
248 2017a) is critical in governing the nitrate variations in Nanjing. During nighttime,
249 the heterogeneous hydrolysis of N_2O_5 can also contribute to nitrate production. It is
250 worth to mention that the nitrate formation mechanisms may vary greatly, for example,
251 during wintertime in Beijing, photochemical production of nitrate is more evident
252 (Sun et al., 2013b). On the other hand, ammonium sulfate is non-volatile and
253 thermodynamic partitioning affects its concentrations in a much lesser extent,
254 however water as a reaction medium, may facilitate the oxidation of SO_2 into sulfate
255 (Seinfeld and Pandis, 2016). In a word, although high RH can promote formations of
256 both nitrate and sulfate, the underlying detailed processes are different.

257 To further investigate the sulfate and nitrate formation mechanisms, we calculated
258 the sulfur oxidation ratio ($SOR = nSO_4^{2-}/(nSO_4^{2-} + nSO_2)$) and nitrogen oxidation ratio
259 ($NOR = nNO_3^-/(nNO_3^- + nNO_2)$). Here nSO_4^{2-} , nNO_3^- , nSO_2 and nNO_2 are the molar
260 concentrations of particle-phase sulfate, nitrate, gaseous sulfur dioxide and nitrogen
261 dioxide, respectively. Note since there are multiple gas-phase forms of nitrogen
262 oxides (NO , N_2O_3 , N_2O_4 , N_2O_5 , etc) rather than only NO_2 , the NOR actually should



263 be smaller than those calculated here. The dependences of NOR and SOR on RH are
264 shown in Figs. 5a and 5b. The NOR was elevated from 0.09-0.13 (average: 0.12) at
265 $RH < 65\%$ to 0.19-0.22 (average: 0.21) at $RH > 65\%$, yet there was no significant
266 enhancement when RH increased from 65% to 95%. Different from NOR, SOR
267 increased nearly linearly with RH from 0.07 at $RH = 35\%$ to 0.50 at $RH = 95\%$ by a
268 large factor of 7.4. The mean SOR under $RH > 80\%$ was 0.46, much higher than 0.23
269 observed in wintertime Beijing at $RH = 80\text{--}90\%$ (Sun et al., 2013a), likely indicating a
270 stronger aqueous-phase SO_2 oxidation ability in Nanjing. The comparison between
271 NOR and SOR shown here demonstrates that moisture in fact plays a more important
272 role for sulfate formation than it does for nitrate. These results also suggest that by
273 considering increases of particle-phase sulfate and nitrate alone, may not well reflect
274 the RH effects, as nitrate concentrations actually increased more rapidly than that of
275 sulfate with the increase of RH: From RH of 40 to 90%, nitrate increased from 5.1 to
276 $16.6 \mu\text{g}/\text{m}^3$, while sulfate increased from 6.8 to $13.9 \mu\text{g}/\text{m}^3$. This result likely suggests
277 that a significant portion of nitrate is due to gas-to-particle conversion rather than
278 direct production in aqueous-phase. Previous studies show that the presence of NO_2
279 actually can enhance aqueous-phase sulfate formation under humid/foggy/cloudy
280 conditions in Nanjing (Xie et al., 2015) and Beijing (Cheng et al., 2016; Wang et al.,
281 2016a), though the significance of this mechanism is heavily dependent upon the RH
282 conditions (Liu et al., 2017; Guo et al., 2017; Zhang et al., 2018b).

283 Correspondingly, Figures 5c and 5d illustrate the variations of NOR and SOR as a
284 function of O_x concentrations. NOR showed no obvious dependence on O_x , suggesting
285 that photochemical production of nitrate is not significant. While SOR displayed an
286 overall decreasing trend with the increase of O_x concentrations. As high RH was
287 associated with low solar radiation (Fig. 1) and low O_x concentrations (Figs. 4a and
288 4b), this result, on the other hand, underscores the dominance of aqueous-phase
289 processing over photochemical processing for sulfate production. In addition, both
290 NOR and SOR did not vary significantly with changes of temperatures (Fig. S3b),
291 again supporting that they were not significantly influenced by photochemical



292 processes. Moreover, the aqueous-phase sulfate production is actually insignificant
293 during springtime in urban Nanjing when the temperature is warmer and the air is
294 drier, as revealed in our previous study (Wang et al., 2016b). Therefore, the findings
295 point out that aerosol sulfate chemistry can be very different under different seasons
296 in the same region. In addition, a measurement study conducted almost in the same
297 period as this study in Beijing (Zhang et al., 2018c) reported the dominance of
298 aqueous-phase production of sulfate, while nitrate was mainly produced by
299 photochemical and heterogeneous reactions. These results indicate that the formation
300 mechanisms of sulfate and nitrate can be different at different locations in the same
301 season.

302

303 3.3 Secondary organic aerosol (SOA) formation

304 3.3.1 Industry-related OA

305 A specific OA factor – industry-related OA (IOA) was resolved by the PMF
306 analysis in this study. This factor on average occupied $\sim 1/6$ of the total OA mass,
307 equivalent to $\sim 1.9 \mu\text{g}/\text{m}^3$. It was different from those from traffic and cooking, as their
308 time series were significantly different (r^2 of 0.03 for HOA vs. IOA, and r^2 of 0.01 for
309 COA vs. IOA). Also, across the sampling period, its temporal variations were
310 relatively smaller than those of HOA and COA, indicating it was a persistent source in
311 this region.

312 The ion-specified HRMS by six ion families is presented in Fig. 6. The IOA O/C
313 ratio (0.44 from I-A method, and 0.33 from A-A method) was relatively high as a
314 POA factor, but was still within a reasonable range. The most abundant ion category
315 in IOA was the chemically-reduced hydrocarbon ions (46%, Fig. 6), but it also had a
316 high contribution from oxygenated ions (especially CO_2^+ and $\text{C}_2\text{H}_3\text{O}^+$). Previously, an
317 industry-related OA factor from the PMF analyses of both online measurement data
318 (El Haddad et al., 2013) and offline $\text{PM}_{2.5}$ samples in Marseille (France) (Bozzetti et
319 al., 2017) was reported. Similar to our IOA, the factor resolved by Bozzetti et al.
320 (2017) had a high CO_2^+ peak ($\sim 10\%$ of total) and a O/C ratio of 0.33. This



321 industry-related OA factor in Marseille was attributed to contributions from coke,
322 steel and petrochemical facilities, which is also very similar to the industry zone
323 adjacent to our site.

324 It should be noted that, the IOA MS had also relatively higher fractions $C_2H_4O_2^+$
325 (m/z 60) and $C_3H_5O_2^+$ (m/z 73) than other factors did. These two ions are often used as
326 biomass burning OA tracers as they can be produced by levoglucosan (Aiken et al.,
327 2010). However, a recent study (Yan et al., 2018) showed that coal combustion could
328 be a significant source of levoglucosan in China. The AMS mass spectrum of peat
329 burning OA was also found to contain appreciable m/z 60 and m/z 73 signals (Lin et
330 al., 2017). Considering that the winds during this campaign were mostly blew from
331 the industry zone with coal (power plants and steel works) and oil burning (petroleum
332 industry) (Fig. S2), and the open biomass burning was banned by the government in
333 this region, the $C_2H_4O_2^+$ and $C_3H_5O_2^+$ signals were likely associated with industry
334 emissions. Of course, more studies are required to further elaborate the characteristics
335 of IOA, including molecular characterization of specific industry-related organic
336 tracer compounds and measurement of the heavy metals which are typically
337 associated with industry emissions.

338 3.3.2 A modified method to separate SOA from POA

339 Previously, Ng et al. (2010) proposed to use a triangle plot of f_{44} (ratio of m/z 44
340 to total signal of the OA factor) versus f_{43} (ratio of m/z 43 to total signal of the OA
341 factor) as a diagnostic to describe different OA factors. Ambient OA in this $f_{44}:f_{43}$
342 space typically converges to the upper left corner with the ageing of OA. This is
343 because f_{44} (mainly CO_2^+) is representative of highly oxygenated carboxylic acids
344 while f_{43} (mainly $C_2H_3O^+$) is mainly produced from non-carboxylic oxygenated
345 species. This plot is presented in Fig. 7b. For this dataset, separation of the factors
346 (especially HOA, IOA, LSOA and SVOOA) was not very well by using this method.
347 Chhabra et al. (2011) also shows that the movement paths of different SOA
348 components in the $f_{44}:f_{43}$ space depend on the precursors. Generally speaking, f_{44}
349 and f_{43} use only two fragments, they may sometimes not well represent the general



350 properties of OA. In addition, m/z 44 and m/z 43 may be significantly influenced by
351 other ions, such as $C_3H_7^+$ on m/z 43 and $C_2H_4O^+/C_2H_6N^+$ on m/z 44.

352 Heringa et al. (2012) proposed another method to discriminate different OA
353 factors. Following this method, we first re-sorted the HRMS according to the number
354 of carbon (C_1 - C_7) and oxygenation state (C_x , C_xO_1 , $C_xO_{>1}$) of each fragment ion. Each
355 HRMS was grouped into 21 clusters and displayed in Fig. 7a. Then, we placed our six
356 PMF-resolved OA factors into the two-dimensional space of C_2/C_2O_1 vs. $C_3O_1/C_3O_{>1}$
357 as suggested by Heringa et al. (2012). As shown in Fig. 7c, the HOA and COA fell far
358 from SVOOA and LVOOA, while IOA and LSOA were still very close to each other.
359 Since the method in fact still uses only a few ions, it is likely better to represent the
360 average properties of OA by making use of the chemical information from more ions.
361 For this purpose, the six OA factors were projected into the fC_xO_y/fC_xR (or
362 $fC_xRO_{>1}/fC_xRO_1$) vs. $fC_{x\geq 3}R/fC_{x\leq 2}R$ spaces (Figs. 7d and 7e). Here fC_xO_y/fC_xR is the
363 ratio of all oxygen-containing ions (C_xO_y) to all oxygen-free ions (C_xR),
364 $fC_xRO_{>1}/fC_xRO_1$ is the ratio of ions with more than one oxygen atoms to those with
365 only one oxygen atom, and $fC_{x\geq 3}R/fC_{x\leq 2}R$ is the ratio of hydrocarbon ions at $x\geq 3$
366 those at $x\leq 2$. As shown in Fig. 7a, with increase of oxidation degrees (O/C) of the OA
367 factors, the contribution from C_x groups at $x\leq 2$ gradually increased, while
368 contribution from hydrocarbon ions (C_xR) with $x\geq 3$ decreased. The small ion
369 fragments also contained more oxygenated ions, likely reflecting the abundance of
370 small organic acids. This trend suggests that the ageing of OA proceeds with
371 fragmentation of large molecules, and a $fC_{x\geq 3}R/fC_{x\leq 2}R$ ratio can reflect such reaction
372 pathway. On the other hand, the ageing of OA often involves the oxidation of
373 hydrocarbon species, thus the fC_xO_y/fC_xR ratio or $fC_xRO_{>1}/fC_xRO_1$ ratio may indicate
374 the oxidation of OA. In both Fig. 7d and Fig. 7e, the SOA factors locate in the upper
375 left region, and can be better separated from the POA factors than that by using the
376 $f_{44}:f_{43}$ space. In particular, the different OA factors approached to the upper left
377 corner in a consecutive order which was consistent with their O/C ratios, and the
378 movement trajectory was also close to a straight line ($y=1.89-0.39x$, $r^2=0.85$). This



379 result suggests that the $(f_{C_xO_y}/f_{C_xR})$: $(f_{C_{x \geq 3}R}/C_{x \leq 2}R)$ space is probably a more
380 effective metric to separate different OA factors and demonstrate the evolution of OA.

381 It should be mentioned that Heald et al. (2010) proposed to use a Van Krevelen
382 diagram (H/C versus O/C) to describe the OA ageing, as shown in Fig. S4 for this
383 dataset. Compared to the f_{44} : f_{43} space (Ng et al., 2010), this method uses the bulk
384 composition of OA. Our modified method proposed here seems to balance the level of
385 chemical information, not only dependent upon a couple of specific ions, but also not
386 rely on the general properties without incorporating detailed ion compositions. Of
387 course, more tests by using more ambient datasets are strongly needed in the future to
388 verify this graphical method.

389 3.3.3 Effects of aqueous-phase and photochemical processing

390 In this study, we identified three SOA factors (LSOA, SVOOA and LVOOA),
391 which together occupied nearly half of OA mass. We plotted the mass concentrations
392 and fractional contributions of the three SOA factors against RH and O_x , to investigate
393 the effects of aqueous-phase and photochemical processing on their formations (Fig.
394 8). Generally, the concentrations of LVOOA presented an increasing trend with RH,
395 and its mass fraction increased more obviously from only 4% at $RH < 40\%$ to 32% at
396 $RH > 90\%$, while those of LSOA and SVOOA showed no such trends. This result
397 demonstrates the importance of aqueous-phase processing on LVOOA formation. On
398 the contrary, both mass loadings and fractions of LSOA and SVOOA increased
399 obviously with the increase of O_x , while those of LVOOA decreased significantly
400 from 46% at $O_x < 50 \mu\text{g}/\text{m}^3$ to 9% at $O_x > 140 \mu\text{g}/\text{m}^3$. Such relationships reveal that
401 photochemical processing promotes the LSOA and SVOOA formations, but
402 contributes negligibly or even hinders the formation of LVOOA. Note the LVOOA
403 had the highest O/C ratio (0.74) among the three SOA factors. This is consistent with
404 the facts that aqSOA typically has a high oxidation level as discovered by a number of
405 lab studies (e.g., Herrmann et al., 2015), and the aqSOA products have low volatility
406 in general and contribute significantly to LVOOA (Ervens et al., 2011). Our results
407 clearly demonstrate the close link between LVOOA and aqSOA. The LSOA and



408 SVOOA were relatively fresher with O/C ratios of 0.44 and 0.66, respectively. Our
409 results suggest that the relatively fresh SOA is closely associated with gasSOA (SOA
410 produced from gas-phase photochemical reactions). These findings are in good
411 agreement with those observed in urban Beijing (Xu et al., 2017a), likely representing
412 the general characteristics of the AMS-resolved OOA factors.

413 For reference, we also checked the variations of POA factors (HOA, COA and
414 IOA) against RH and O_x (Fig. S5). Results show that both concentrations and
415 fractions of the POA factors overall had no clear responses to either RH or O_x
416 increases, which are in accordance with their behaviors as POA – they were emitted
417 directly into the air thus were not expected to be influenced by both aqueous-phase or
418 photochemical processing. One exception was that COA concentrations seemed to
419 increase with O_x , which was likely a coincidence as the lunchtime was exactly
420 noon/early afternoon with strong photochemical activities.

421 To further investigate the influences of aqueous-phase and photochemical
422 processing on the oxidation levels of SOA (O/C_{SOA}), we calculated the O/C_{SOA} based
423 on the method proposed by Xu et al. (2017b). It represents the combined O/C ratio of
424 the summed LSOA, SVOOA and LVOOA. The changes of O/C_{SOA} and O/C_{COA} (the
425 O/C ratios of total OA) versus RH and O_x were shown in Figs. 8e and 8f. The O/C_{COA}
426 showed no clear dependences on both RH and O_x , probably due to the mixing effects
427 of POA. However, the O/C_{SOA} tended to be larger at higher RH (especially at
428 $RH > 60\%$), indicating the contribution of more oxygenated aqSOA; while the O/C_{SOA}
429 clearly decreased with increase of O_x concentrations (from 0.71 at $O_x < 50 \mu\text{g}/\text{m}^3$ to
430 0.57 at $O_x > 140 \mu\text{g}/\text{m}^3$), as more photochemical SOA with lower O/C ratios were
431 generated.

432 The combined effects of RH and O_x were further illustrated in Fig. 9. High
433 SVOOA and LSOA concentrations appeared in the regimes with high O_x
434 concentrations (Figs. 9a and 9b) and high fractions of these two factors also tended to
435 be accompanied more significantly with low RH conditions (Figs. 9d and 9e).
436 Differently, both high concentrations and high fractions of LVOOA located in the



437 upper left corner characterized by high RH and low O_x conditions (Figs. 9c and 9f),
438 again supporting the importance of aqueous-phase processing on LVOOA production.

439 Furthermore, it is known that different formation pathways have different impacts
440 on the size distributions of fine aerosols. Secondary species from gas-phase
441 photochemical reactions typically condense on smaller particles (condensation mode),
442 while particles from aqueous-phase reactions in clouds/fogs or wet aerosols are larger
443 (droplet mode) (Ervens et al., 2011; Kerminen and Wexler, 1995; Meng and Seinfeld,
444 1994). Enhancements of particle sizes were indeed observed when aqueous-phase
445 reactions occurred (Gilarioni et al., 2016; Ge et al., 2012). In this regard, we
446 investigated the average size distributions of OA, sulfate and nitrate at different RH
447 levels (Fig. 10). Obviously, the peak sizes of OA, sulfate and nitrate all shifted
448 towards larger sizes with the increase of RH. The OA size distribution peaked at 268
449 nm *D*_{va} (vacuum aerodynamic diameter) at RH<40%, and increased substantially to
450 694 nm *D*_{va} at RH of 88-92%. Because increase of RH did not elevate the three POA
451 (Fig. S5), LSOA and SVOOA factors (Figs. 8a and 8b), the growth of OA sizes can be
452 attributed mostly to the formation of LVOOA, namely aqSOA. The peak *D*_{va} of sulfate
453 increased from 418 nm (RH<40%) to 730 nm (RH=88-92%), and that of nitrate
454 increased even more significantly – from 233 nm (RH<40%) to 694 nm
455 (RH=88-92%). Small drops of the peak *D*_{va} occurring at RH>92% for all three species,
456 were probably due to the scavenging effects of precipitation. Overall, the
457 enhancements of peak sizes of OA, sulfate and nitrate by RH provide another
458 evidence supporting the importance of aqueous-phase processing on secondary
459 aerosol formation in this study.

460

461 **3.4. Influences of secondary aerosol formation on PM₁ pollution**

462 **3.4.1 Mass contributions at different pollution levels**

463 We have demonstrated that moisture plays important roles to the formations of
464 sulfate, nitrate and LVOOA, while photochemical processing is important for LSOA
465 and SVOOA productions. Here we investigated contributions of these secondary



466 components at different levels of PM₁ and OA pollutions. As shown in Fig. 11a, mass
467 fractions of nitrate increased from 16.5% during clean periods (PM₁<10 μg/m³) to
468 ~26-27 % during polluted periods (PM₁>60 μg/m³) (there was a small drop from 27.1%
469 at PM₁ of 60-70 μg/m³ to 25.9% at PM₁>70 μg/m³). Variations of sulfate contribution
470 were relatively small from 25.2% at PM₁<10 μg/m³ to 21.1% at PM₁>60 μg/m³. We
471 also illustrated the mean RH and O_x concentrations at different PM₁ levels in Fig. 11a.
472 With the aggravation of PM₁ pollution, RH also significantly increased (from 57% to
473 83%), but O_x concentrations varied very little within a range of 82.3 - 90.3 μg/m³.
474 This result shows the importance of moisture to haze pollution as it promotes the
475 formations of nitrate and sulfate. In fact, both NOR and SOR values increased with
476 the increase of PM₁ concentrations (Fig. S3a). As we showed earlier, nitrate
477 concentrations increased more quickly with RH than that of sulfate, thus its fractions
478 increased significantly, while the sulfate concentrations increased but its relative
479 contributions decreased slightly. The faster increase of nitrate concentrations than
480 sulfate during heavy pollution was also observed in Nanjing during 2014 (Wu et al.,
481 2017), Nanjing (Zhang et al., 2015b) and Lin'an (Shen et al., 2015) during 2013
482 January in the YRD region. This is different from that in North China Plain, where
483 sulfate often plays a more significant role in heavy haze formation.

484 For organics, the contributions continuously decreased from 32.3% at PM₁<10
485 μg/m³ to 23.6% at PM₁>70 μg/m³ (Fig. 11a). This result indicates that although high
486 RH could enhance the production of a portion of SOA (namely the LVOOA), overall,
487 the OA pollution was not governed by RH effects. Therefore, we calculated the mass
488 fractions of different OA factors to total OA at different OA levels in Fig. 11b. Clearly,
489 the LVOOA contributions changed from 25.2-29.0% at OA<10 μg/m³ down to only
490 10.4% at OA>30 μg/m³, while the sum of LSOA and SVOOA increased substantially
491 from 12.1% at OA<5 μg/m³ to 45.1% at OA>30 μg/m³. This result shows that
492 photochemical processing was more important than the aqueous-phase processing in
493 exacerbating the OA pollution. Correspondingly, we found that the RH did not vary
494 significantly as a function of OA concentrations (within 70-80%), while O_x



495 concentrations increased obviously from $77.6 \mu\text{g}/\text{m}^3$ at $\text{OA} < 5 \mu\text{g}/\text{m}^3$ to $115.7 \mu\text{g}/\text{m}^3$ at
496 $\text{OA} > 30 \mu\text{g}/\text{m}^3$. This is somewhat opposite to the case of PM_{10} , and partially explains
497 the decrease of OA contributions to the heavy PM_{10} pollutions. In addition, the total
498 POA mass fractions decreased from 62.7% at $\text{OA} < 5 \mu\text{g}/\text{m}^3$ to 44.5% at $\text{OA} > 30 \mu\text{g}/\text{m}^3$.
499 This is different from that found during springtime in urban Nanjing (Wang et al.,
500 2016b), where the heavier OA pollution periods were accompanied by elevated POA
501 contributions.

502 Moreover, we calculated $f\text{CO}_2^+/f\text{C}_2\text{H}_3\text{O}^+$ and Y/X ($Y=f\text{C}_x\text{O}_y/f\text{C}_x\text{R}$, $X=f\text{C}_{x\geq 3}\text{R}/f\text{C}_{x\leq 2}\text{R}$)
503 against OA concentrations in Fig. 11c. The $f\text{CO}_2^+/f\text{C}_2\text{H}_3\text{O}^+$ modified the f_{44}/f_{43}
504 index (Ng et al., 2010) as it eliminates influences from other ions on m/z 44 and m/z
505 43. Both indexes decreased with OA concentrations, showing that the reduction of
506 highly oxygenated LVOOA contributions and increase of moderately oxidized LSOA
507 and SVOOA contributions compensated the decrease of fresh POA contributions.
508 Correspondingly, $f\text{C}_4\text{H}_9^+$ and $f\text{C}_{x\geq 3}\text{R}/f\text{C}_{x\leq 2}\text{R}$ increased with the increase of OA
509 concentrations. As C_4H_9^+ (m/z 57) is often used a POA tracer, and $f\text{C}_{x\geq 3}\text{R}/f\text{C}_{x\leq 2}\text{R}$ also
510 reflects the abundance of chemically reduced hydrocarbon ions (Section 3.3.2), this
511 result again suggests the OA overall became less oxygenated when its pollution
512 became heavier. All results underscore that the relatively fresh and photo-formed SOA
513 (LSOA and SVOOA) plays dominant roles to the OA pollution in this industrialized
514 environment.

515 In addition, we checked the variations of mass fractions of different components
516 at different wind speeds (WS) and wind directions (WD) (Fig. S6). Generally, the
517 changes were not dramatic, consistent with those of RH (67%-79%) and O_x (79-95
518 $\mu\text{g}/\text{m}^3$), which did not present clear increasing or decreasing trends against WS and
519 WD as well. Such results indicate that the influences of secondary aerosol species
520 seemed to be not affected significantly by the different air masses in this location.

521 3.4.2 Case studies

522 We investigated influences of secondary aerosol formation in two specific cases.
523 The first case was from 3:30pm March 1 to 3:30 pm March 2 (marked in Fig. 2).



524 During this episode, the mean PM₁ concentrations was 48.9 μg/m³, and the summed
525 mass contributions from (NH₄)₂SO₄, NH₄NO₃ and LVOOA to PM₁ was 74.4%, both
526 higher than their corresponding campaign-average values (46.3 μg/m³ and 68.6%, Fig.
527 3a). This episode was characterized by significant productions of sulfate, nitrate and
528 LVOOA (SNL). The total PM₁ concentrations were very closely linked with the SNL
529 formation (Fig. 12b) especially during later afternoon and nighttime. Correspondingly,
530 the enhancement of SNL almost linearly responded to the increase of RH ($r=0.96$),
531 but oppositely correlated with that of O_x ($r=-0.86$). High RH during nighttime were
532 associated with high SNL concentrations, while high O_x concentrations during
533 noon/early afternoon were accompanied by low SNL concentrations.

534 Another case was characterized by relatively significant photochemical formation
535 of SOA, especially the LSOA. This episode lasted from the early evening of March 12
536 till early morning of March 16 (marked in Fig. 2). In fact, high LSOA concentrations
537 mainly occurred in this period, which was on average 4.8 μg/m³, much higher than its
538 average value of 0.38 μg/m³ during other periods. This is actually one reason it was
539 defined as LSOA since it was most likely a specific and localized event. We found it
540 was mainly driven by photochemical processing (Section 3.3.3), but unable to identify
541 the precursors whose emissions were particularly enhanced during this episode and
542 led to the LSOA formation due to measurement limitations. During this period, the
543 mean OA concentration was 18.3 μg/m³ (Fig. 13a), much higher than its
544 campaign-average value of 11.8 μg/m³; mass fraction of LSOA and SVOOA (LSS)
545 together was 43.6% (Fig. 13a), also higher than the campaign average of 26.4%.
546 Similar as those in Fig. 8, we observed the enhancements of LSS mass concentrations
547 and fractions at high O_x concentrations but reductions at high RH conditions (Fig. S7).
548 More clearly, from noontime of March 15 to early morning of March 16, the LSS
549 mass concentrations correlated positively very well with O_x ($r=0.91$) but meanwhile
550 linearly decreased with RH ($r=-0.90$) (Figs. 13b and 13c). Contrary to the first case
551 (Fig. 12), high O_x concentrations during noon/afternoon associated with high LSS
552 concentrations, while high RH during nighttime associated with low LSS



553 concentrations. These two cases are two typical examples demonstrating especially
554 the aqueous-phase nighttime formation of SNL and the daytime photo-formation of
555 LSS, respectively.

556

557 **4. Conclusions**

558 This work presents the field measurement results focusing on the formation and
559 characteristics of secondary inorganic (mainly sulfate, nitrate) and organic species, by
560 using the SP-AMS in suburban Nanjing during February – March, 2015. The site was
561 surrounded by a large number of industrial plants. Nevertheless, under this
562 industrialized environment, the PM₁ was mainly comprised of secondary aerosol
563 components (75.4%), including 63.2% from ammonium sulfate and nitrate, and 12.2%
564 from SOA. This finding indicates that the gas-phase precursors such as SO₂, NO₂ and
565 volatile organic compounds can be effectively transformed into particle-phase species.
566 Furthermore, moisture was found to play a major role in enhancing productions of
567 sulfate, nitrate and the highly oxygenated portion of SOA (LVOOA, 45% of SOA),
568 yet the detailed mechanisms were different. The moisture likely affected nitrate by
569 enhancing its thermodynamic gas-particle partitioning and nocturnal heterogeneous
570 production, while direct productions of sulfate and LVOOA in aqueous phase were
571 more significant. In addition, the peak sizes of sulfate, nitrate and OA all shifted
572 towards larger sizes with the increases of relative humidity, reflecting the effects of
573 aqueous-phase processing too. On the other hand, the other two less oxygenated SOA
574 factors (LSOA and SVOOA, together 55% of SOA) were mainly driven by
575 photochemical processing.

576 Overall, the moisture-driven nitrate and sulfate productions were important to
577 aggravate the PM₁ pollution. The photo-chemically formed SOA was more important
578 than the aqueous-phase SOA to OA pollution. The influences of these two formation
579 pathways were demonstrated very clearly in two typical episodes. In addition, in this
580 study we also provided two new findings. First, we separated a specific
581 industry-related OA factor. Secondly, we proposed a modified graphical method to



582 describe the evolution of OA. Both of them should be investigated and verified in the
583 future with other AMS datasets. In summary, this work highlights that aqueous-phase
584 chemistry is very important for sulfate and nitrate productions in an industrialized
585 environment during cold seasons. We also show that the highly oxygenated SOA is
586 very likely linked with aqueous-phase processing while the less oxygenated ones are
587 associated with photochemical processing.

588

589 **5. Data availability**

590 The observational data in this study are available from the authors upon request
591 (caxinra@163.com).

592

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601 **References**

602 Aiken, A. C., Decarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty, K. S., Ulbrich, I.
603 M., Mohr, C., Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q., Trimborn, A., Northway, M., Ziemann, P.
604 J., Canagaratna, M. R., Onasch, T. B., Alfarra, M. R., Prevot, A. S. H., Dommen, J., Duplissy, J.,
605 Metzger, A., Baltensperger, U., and Jimenez, J. L.: O/C and OM/OC ratios of primary, secondary, and
606 ambient organic aerosols with high-resolution time-of-flight aerosol mass spectrometry, *Environ. Sci.*
607 *Technol.*, 42, 4478-4485, 10.1021/Es703009q, 2008.

608 Aiken, A. C., de Foy, B., Wiedinmyer, C., DeCarlo, P. F., Ulbrich, I. M., Wehrli, M. N., Szidat, S.,
609 Prevot, A. S. H., Noda, J., Wacker, L., Volkamer, R., Fortner, E., Wang, J., Laskin, A., Shutthanandan,
610 V., Zheng, J., Zhang, R., Paredes-Miranda, G., Arnott, W. P., Molina, L. T., Sosa, G., Querol, X., and
611 Jimenez, J. L.: Mexico city aerosol analysis during MILAGRO using high resolution aerosol mass
612 spectrometry at the urban supersite (T0) - Part 2: Analysis of the biomass burning contribution and the
613 non-fossil carbon fraction, *Atmos. Chem. Phys.*, 10, 5315-5341, 10.5194/acp-10-5315-2010, 2010.

614 Blando, J. D., and Turpin, B. J.: Secondary organic aerosol formation in cloud and fog droplets: a



- 615 literature evaluation of plausibility, *Atmos. Environ.*, 34, 1623-1632, 10.1016/s1352-2310(99)00392-1,
616 2000.
- 617 Bozzetti, C., El Haddad, I., Salameh, D., Daellenbach, K. R., Fermo, P., Gonzalez, R., Minguillón, M.
618 C., Iinuma, Y., Poulain, L., Elser, M., Müller, E., Slowik, J. G., Jaffrezo, J. L., Baltensperger, U.,
619 Marchand, N., and Prévôt, A. S. H.: Organic aerosol source apportionment by offline-AMS over a full
620 year in Marseille, *Atmos. Chem. Phys.*, 17, 8247-8268, 10.5194/acp-17-8247-2017, 2017.
- 621 Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q., Onasch, T. B.,
622 Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R., Trimborn, A. M., Northway, M. J.,
623 DeCarlo, P. F., Kolb, C. E., Davidovits, P., and Worsnop, D. R.: Chemical and microphysical
624 characterization of ambient aerosols with the aerodyne aerosol mass spectrometer, *Mass Spectrom.*
625 *Rev.*, 26, 185-222, 10.1002/Mas.20115, 2007.
- 626 Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P., Hildebrandt Ruiz,
627 L., Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue, N. M., Jayne, J. T., and Worsnop,
628 D. R.: Elemental ratio measurements of organic compounds using aerosol mass spectrometry:
629 characterization, improved calibration, and implications, *Atmos. Chem. Phys.*, 15, 253-272,
630 10.5194/acp-15-253-2015, 2015.
- 631 Carlton, A. G., Turpin, B. J., Lim, H. J., Altieri, K. E., and Seitzinger, S.: Link between isoprene and
632 secondary organic aerosol (SOA): Pyruvic acid oxidation yields low volatility organic acids in clouds,
633 *Geophys. Res. Lett.*, 33, 10.1029/2005gl025374, 2006.
- 634 Carslaw, K. S., Boucher, O., Spracklen, D. V., Mann, G. W., Rae, J. G. L., Woodward, S., and Kulmala,
635 M.: A review of natural aerosol interactions and feedbacks within the Earth system, *Atmos. Chem.*
636 *Phys.*, 10, 1701-1737, 10.5194/acp-10-1701-2010, 2010.
- 637 Chan, C. K., and Yao, X.: Air pollution in mega cities in China, *Atmos. Environ.*, 42, 1-42,
638 10.1016/j.atmosenv.2007.09.003, 2008.
- 639 Cheng, Y., Zheng, G., Wei, C., Mu, Q., Zheng, B., Wang, Z., Gao, M., Zhang, Q., He, K., Carmichael,
640 G., Pöschl, U., and Su, H.: Reactive nitrogen chemistry in aerosol water as a source of sulfate during
641 haze events in China, *Sci. Adv.*, 2, e1601530, 10.1126/sciadv.1601530, 2016.
- 642 Chhabra, P. S., Ng, N. L., Canagaratna, M. R., Corrigan, A. L., Russell, L. M., Worsnop, D. R., Flagan,
643 R. C., and Seinfeld, J. H.: Elemental composition and oxidation of chamber organic aerosol, *Atmos.*
644 *Chem. Phys.*, 11, 8827-8845, 10.5194/acp-11-8827-2011, 2011.
- 645 De Haan, D. O., Hawkins, L. N., Kononenko, J. A., Turley, J. J., Corrigan, A. L., Tolbert, M. A., and
646 Jimenez, J. L.: Formation of nitrogen-nontaining oligomers by methylglyoxal and amines in simulated
647 evaporating cloud droplets, *Environ. Sci. Technol.*, 45, 984-991, 10.1021/es102933x, 2011.
- 648 El Haddad, I., D'Anna, B., Temime-Roussel, B., Nicolas, M., Boreave, A., Favez, O., Voisin, D., Sciare,
649 J., George, C., Jaffrezo, J. L., Wortham, H., and Marchand, N.: Towards a better understanding of the
650 origins, chemical composition and aging of oxygenated organic aerosols: case study of a Mediterranean
651 industrialized environment, Marseille, *Atmos. Chem. Phys.*, 13, 7875-7894, 10.5194/acp-13-7875-2013,



- 652 2013.
- 653 Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in cloud droplets and
654 aqueous particles (aqSOA): a review of laboratory, field and model studies, *Atmos. Chem. Phys.*, 11,
655 11069-11102, 10.5194/acp-11-11069-2011, 2011.
- 656 Ervens, B.: Modeling the processing of aerosol and trace gases in clouds and fogs, *Chem. Rev.*, 115,
657 4157-4198, 10.1021/cr5005887, 2015.
- 658 Ge, X., Wexler, A. S., and Clegg, S. L.: Atmospheric amines - Part II. Thermodynamic properties and
659 gas/particle partitioning, *Atmos. Environ.*, 45, 561-577, 10.1016/j.atmosenv.2010.10.013, 2011a.
- 660 Ge, X., Wexler, A. S., and Clegg, S. L.: Atmospheric amines - Part I. A review, *Atmos. Environ.*, 45,
661 524-546, 10.1016/j.atmosenv.2010.10.012, 2011b.
- 662 Ge, X., Zhang, Q., Sun, Y., Ruehl, C. R., and Setyan, A.: Effect of aqueous-phase processing on aerosol
663 chemistry and size distributions in Fresno, California, during wintertime, *Environ. Chem.*, 9, 221-235,
664 10.1071/EN11168, 2012.
- 665 Ge, X., He, Y., Sun, Y., Xu, J., Wang, J., Shen, Y., and Chen, M.: Characteristics and formation
666 mechanisms of fine particulate nitrate in typical urban areas in China, *Atmosphere*, 8, 62, 2017a.
- 667 Ge, X., Li, L., Chen, Y., Chen, H., Wu, D., Wang, J., Xie, X., Ge, S., Ye, Z., Xu, J., and Chen, M.:
668 Aerosol characteristics and sources in Yangzhou, China resolved by offline aerosol mass spectrometry
669 and other techniques, *Environ. Pollut.*, 225, 74-85, 10.1016/j.envpol.2017.03.044, 2017b.
- 670 George, C., Ammann, M., D'Anna, B., Donaldson, D. J., and Nizkorodov, S. A.: Heterogeneous
671 photochemistry in the atmosphere, *Chem. Rev.*, 115, 4218-4258, 10.1021/cr500648z, 2015a.
- 672 George, K. M., Ruthenburg, T. C., Smith, J., Yu, L., Zhang, Q., Anastasio, C., and Dillner, A. M.: FT-IR
673 quantification of the carbonyl functional group in aqueous-phase secondary organic aerosol from
674 phenols, *Atmos. Environ.*, 100, 230-237, 10.1016/j.atmosenv.2014.11.011, 2015b.
- 675 Gilardoni, S., Massoli, P., Paglione, M., Giulianelli, L., Carbone, C., Rinaldi, M., Decesari, S., Sandrini,
676 S., Costabile, F., Gobbi, G. P., Pietrogrande, M. C., Visentin, M., Scotto, F., Fuzzi, S., and Facchini, M.
677 C.: Direct observation of aqueous secondary organic aerosol from biomass-burning emissions, *P. Natl.
678 Acad. Sci. USA*, 10.1073/pnas.1602212113, 2016.
- 679 Guo, H., Weber, R. J., and Nenes, A.: High levels of ammonia do not raise fine particle pH sufficiently
680 to yield nitrogen oxide-dominated sulfate production, *Sci. Rep.*, 7, 12109,
681 10.1038/s41598-017-11704-0, 2017.
- 682 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J.,
683 Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y.,
684 Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T.
685 F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The
686 formation, properties and impact of secondary organic aerosol: current and emerging issues, *Atmos.*



- 687 Chem. Phys., 9, 5155-5236, 10.5194/acp-9-5155-2009, 2009.
- 688 Heal, M. R., Kumar, P., and Harrison, R. M.: Particles, air quality, policy and health, Chem. Soc. Rev.,
689 41, 6606-6630, 2012.
- 690 Heald, C. L., Jacob, D. J., Park, R. J., Russell, L. M., Huebert, B. J., Seinfeld, J. H., Liao, H., and
691 Weber, R. J.: A large organic aerosol source in the free troposphere missing from current models,
692 Geophys. Res. Lett., 32, L18809, 10.1029/2005GL023831, 2005.
- 693 Heald, C. L., Kroll, J. H., Jimenez, J. L., Docherty, K. S., DeCarlo, P. F., Aiken, A. C., Chen, Q., Martin,
694 S. T., Farmer, D. K., and Artaxo, P.: A simplified description of the evolution of organic aerosol
695 composition in the atmosphere, Geophys. Res. Lett., 37, L08803, 10.1029/2010gl042737, 2010.
- 696 Heringa, M. F., DeCarlo, P. F., Chirico, R., Tritscher, T., Clairotte, M., Mohr, C., Crippa, M., Slowik, J.
697 G., Pfaffenberger, L., Dommen, J., Weingartner, E., Prévôt, A. S. H., and Baltensperger, U.: A new
698 method to discriminate secondary organic aerosols from different sources using high-resolution aerosol
699 mass spectra, Atmos. Chem. Phys., 12, 2189-2203, 10.5194/acp-12-2189-2012, 2012.
- 700 Herndon, S. C., Onasch, T. B., Wood, E. C., Kroll, J. H., Canagaratna, M. R., Jayne, J. T., Zavala, M. A.,
701 Knighton, W. B., Mazzoleni, C., Dubey, M. K., Ulbrich, I. M., Jimenez, J. L., Seila, R., de Gouw, J. A.,
702 de Foy, B., Fast, J., Molina, L. T., Kolb, C. E., and Worsnop, D. R.: Correlation of secondary organic
703 aerosol with odd oxygen in Mexico City, Geophys. Res. Lett., 35, L15804, 10.1029/2008gl034058,
704 2008.
- 705 Herrmann, H., Schaefer, T., Tilgner, A., Styler, S. A., Weller, C., Teich, M., and Otto, T.: Tropospheric
706 aqueous-phase chemistry: Kinetics, mechanisms, and its coupling to a changing gas phase, Chem. Rev.,
707 115, 4259-4334, 10.1021/cr500447k, 2015.
- 708 Hu, J., Wang, Y., Ying, Q., and Zhang, H.: Spatial and temporal variability of PM_{2.5} and PM₁₀ over
709 the North China Plain and the Yangtze River Delta, China, Atmos. Environ., 95, 598-609,
710 10.1016/j.atmosenv.2014.07.019, 2014.
- 711 Hu, J., Huang, L., Chen, M., Liao, H., Zhang, H., Wang, S., Zhang, Q., and Ying, Q.: Premature
712 mortality attributable to particulate matter in China: Source contributions and responses to reductions,
713 Environ. Sci. Technol., 51, 9950-9959, 10.1021/acs.est.7b03193, 2017.
- 714 Huang, R., Zhang, Y., Bozzetti, C., Ho, K., Cao, J., Han, Y., Daellenbach, K. R., Slowik, J. G., Platt, S.
715 M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa, M., Ciarelli, G., Piazzalunga,
716 A., Schwikowski, M., Abbaszade, G., Schnelle-Kreis, J., Zimmermann, R., An, Z., Szidat, S.,
717 Baltensperger, U., Haddad, I. E., and Prevot, A. S. H.: High secondary aerosol contribution to
718 particulate pollution during haze events in China, Nature, 514, 218-222, 10.1038/nature13774, 2014.
- 719 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo,
720 P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P.,
721 Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J.,
722 Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M.,
723 Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P.



- 724 I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K.,
725 Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y.,
726 Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M.,
727 Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.:
728 Evolution of organic aerosols in the atmosphere, *Science*, 326, 1525-1529, 10.1126/science.1180353,
729 2009.
- 730 Kerminen, V.-M., and Wexler, A. S.: Growth laws for atmospheric aerosol particles: An examination of
731 the bimodality of the accumulation mode, *Atmos. Environ.*, 29, 3263-3275,
732 10.1016/1352-2310(95)00249-X, 1995.
- 733 Li, K., Chen, L., White, S. J., Zheng, X., Lv, B., Lin, C., Bao, Z., Wu, X., Gao, X., Ying, F., Shen, J.,
734 Azzi, M., and Cen, K.: Chemical characteristics and sources of PM₁ during the 2016 summer in
735 Hangzhou, *Environ. Pollut.*, 232, 42-54, 10.1016/j.envpol.2017.09.016, 2018.
- 736 Li, Y., Sun, Y., Zhang, Q., Li, X., Li, M., Zhou, Z., and Chan, C. K.: Real-time chemical
737 characterization of atmospheric particulate matter in China: A review, *Atmos. Environ.*, 158, 270-304,
738 10.1016/j.atmosenv.2017.02.027, 2017.
- 739 Li, Y. J., Lee, B. Y. L., Yu, J. Z., Ng, N. L., and Chan, C. K.: Evaluating the degree of oxygenation of
740 organic aerosol during foggy and hazy days in Hong Kong using high-resolution time-of-flight aerosol
741 mass spectrometry (HR-ToF-AMS), *Atmos. Chem. Phys.*, 13, 8739-8753, 10.5194/acp-13-8739-2013,
742 2013.
- 743 Lim, Y. B., Tan, Y., Perri, M. J., Seitzinger, S. P., and Turpin, B. J.: Aqueous chemistry and its role in
744 secondary organic aerosol (SOA) formation, *Atmos. Chem. Phys.*, 10, 10521-10539,
745 10.5194/acp-10-10521-2010, 2010.
- 746 Lim, Y. B., Tan, Y., and Turpin, B. J.: Chemical insights, explicit chemistry, and yields of secondary
747 organic aerosol from OH radical oxidation of methylglyoxal and glyoxal in the aqueous phase, *Atmos.*
748 *Chem. Phys.*, 13, 8651-8667, 10.5194/acp-13-8651-2013, 2013.
- 749 Lin, C., Ceburnis, D., Hellebust, S., Buckley, P., Wenger, J., Canonaco, F., Prévôt, A. S. H., Huang, R.,
750 O'Dowd, C., and Ovadnevaite, J.: Characterization of Primary Organic Aerosol from Domestic Wood,
751 Peat, and Coal Burning in Ireland, *Environ. Sci. Technol.*, 51, 10624-10632, 10.1021/acs.est.7b01926,
752 2017.
- 753 Liu, M., Song, Y., Zhou, T., Xu, Z., Yan, C., Zheng, M., Wu, Z., Hu, M., Wu, Y., and Zhu, T.: Fine
754 particle pH during severe haze episodes in northern China, *Geophys. Res. Lett.*, 44, 5213-5221,
755 10.1002/2017GL073210, 2017.
- 756 Liu, Y., Monod, A., Tritscher, T., Praplan, A. P., DeCarlo, P. F., Temime-Roussel, B., Quivet, E.,
757 Marchand, N., Dommen, J., and Baltensperger, U.: Aqueous phase processing of secondary organic
758 aerosol from isoprene photooxidation, *Atmos. Chem. Phys.*, 12, 5879-5895, 10.5194/acp-12-5879-2012,
759 2012.
- 760 McNeill, V. F.: Aqueous organic chemistry in the atmosphere: Sources and chemical processing of



- 761 organic aerosols, *Environ. Sci. Technol.*, 49, 1237-1244, 10.1021/es5043707, 2015.
- 762 Meng, Z., and Seinfeld, J. H.: On the source of the submicrometer droplet mode of urban and regional
763 aerosols, *Aerosol Sci. Tech.*, 20, 253-265, 10.1080/02786829408959681, 1994.
- 764 Middlebrook, A. M., Bahreini, R., Jimenez, J. L., and Canagaratna, M. R.: Evaluation of
765 composition-dependent collection efficiencies for the Aerodyne aerosol mass spectrometer using field
766 data, *Aerosol Sci. Tech.*, 46, 258-271, 10.1080/02786826.2011.620041, 2012.
- 767 Ng, N. L., Canagaratna, M. R., Zhang, Q., Jimenez, J. L., Tian, J., Ulbrich, I. M., Kroll, J. H., Docherty,
768 K. S., Chhabra, P. S., Bahreini, R., Murphy, S. M., Seinfeld, J. H., Hildebrandt, L., Donahue, N. M.,
769 DeCarlo, P. F., Lanz, V. A., Prevot, A. S. H., Dinar, E., Rudich, Y., and Worsnop, D. R.: Organic aerosol
770 components observed in Northern Hemispheric datasets from Aerosol Mass Spectrometry, *Atmos.*
771 *Chem. Phys.*, 10, 4625-4641, 10.5194/acp-10-4625-2010, 2010.
- 772 Onasch, T. B., Trimborn, A., Fortner, E. C., Jayne, J. T., Kok, G. L., Williams, L. R., Davidovits, P., and
773 Worsnop, D. R.: Soot particle aerosol mass spectrometer: Development, validation, and initial
774 application, *Aerosol Sci. Tech.*, 46, 804-817, 10.1080/02786826.2012.663948, 2012.
- 775 Pöschl, U.: Atmospheric aerosols: Composition, transformation, climate and health effects, *Angew.*
776 *Chem., Int. Ed.*, 44, 7520-7540, 10.1002/anie.200501122, 2005.
- 777 Paatero, P., and Tapper, U.: Positive matrix factorization: A non-negative factor model with optimal
778 utilization of error estimates of data values, *Environmetrics*, 5, 111-126, 10.1002/env.3170050203,
779 1994.
- 780 Philip, S., Martin, R. V., Pierce, J. R., Jimenez, J. L., Zhang, Q., Canagaratna, M. R., Spracklen, D. V.,
781 Nowlan, C. R., Lamsal, L. N., Cooper, M. J., and Krotkov, N. A.: Spatially and seasonally resolved
782 estimate of the ratio of organic mass to organic carbon, *Atmos. Environ.*, 87, 34-40,
783 10.1016/j.atmosenv.2013.11.065, 2014.
- 784 Pope, C. A., and Dockery, D. W.: Health effects of fine particulate air pollution: Lines that connect, *J.*
785 *Air Waste Manage.*, 56, 709-742, 10.1080/10473289.2006.10464485, 2006.
- 786 Saffari, A., Hasheminassab, S., Shafer, M. M., Schauer, J. J., Chatila, T. A., and Sioutas, C.: Nighttime
787 aqueous-phase secondary organic aerosols in Los Angeles and its implication for fine particulate matter
788 composition and oxidative potential, *Atmos. Environ.*, 133, 112-122, 10.1016/j.atmosenv.2016.03.022,
789 2016.
- 790 Seinfeld, J. H., and Pandis, S. N.: Atmospheric chemistry and physics: From air pollution to climate
791 change, John Wiley & Sons, New York, 2016.
- 792 Shen, F., Ge, X., Hu, J., Nie, D., Tian, L., and Chen, M.: Air pollution characteristics and health risks in
793 Henan Province, China, *Environ. Res.*, 156, 625-634, 10.1016/j.envres.2017.04.026, 2017.
- 794 Shen, G., Xue, M., Yuan, S., Zhang, J., Zhao, Q., Li, B., Wu, H., and Ding, A.: Chemical compositions
795 and reconstructed light extinction coefficients of particulate matter in a mega-city in the western



- 796 Yangtze River Delta, China, Atmos. Environ., 83, 14-20, 10.1016/j.atmosenv.2013.10.055, 2014.
- 797 Shen, X. J., Sun, J. Y., Zhang, X. Y., Zhang, Y. M., Zhang, L., Che, H. C., Ma, Q. L., Yu, X. M., Yue, Y.,
798 and Zhang, Y. W.: Characterization of submicron aerosols and effect on visibility during a severe
799 haze-fog episode in Yangtze River Delta, China, Atmos. Environ., 120, 307-316,
800 10.1016/j.atmosenv.2015.09.011, 2015.
- 801 Shiraiwa, M., Ueda, K., Pozzer, A., Lammel, G., Kampf, C. J., Fushimi, A., Enami, S., Arangio, A. M.,
802 Fröhlich-Nowoisky, J., Fujitani, Y., Furuyama, A., Lakey, P. S. J., Lelieveld, J., Lucas, K., Morino, Y.,
803 Pöschl, U., Takahama, S., Takami, A., Tong, H., Weber, B., Yoshino, A., and Sato, K.: Aerosol health
804 effects from molecular to global scales, Environ. Sci. Technol., 51, 13545-13567,
805 10.1021/acs.est.7b04417, 2017.
- 806 Spracklen, D. V., Jimenez, J. L., Carslaw, K. S., Worsnop, D. R., Evans, M. J., Mann, G. W., Zhang, Q.,
807 Canagaratna, M. R., Allan, J., Coe, H., McFiggans, G., Rap, A., and Forster, P.: Aerosol mass
808 spectrometer constraint on the global secondary organic aerosol budget, Atmos. Chem. Phys., 11,
809 12109-12136, 10.5194/acp-11-12109-2011, 2011.
- 810 ToF-AMS Analysis Software, available at:
811 <http://cires1.colorado.edu/jimenez-group/ToFAMSResources/ToFSoftware/index.html>, 2015.
- 812 Sun, Y., Wang, Z., Fu, P., Jiang, Q., Yang, T., Li, J., and Ge, X.: The impact of relative humidity on
813 aerosol composition and evolution processes during wintertime in Beijing, China, Atmos. Environ., 77,
814 927-934, 10.1016/j.atmosenv.2013.06.019, 2013a.
- 815 Sun, Y., Du, W., Fu, P., Wang, Q., Li, J., Ge, X., Zhang, Q., Zhu, C., Ren, L., Xu, W., Zhao, J., Han, T.,
816 Worsnop, D. R., and Wang, Z.: Primary and secondary aerosols in Beijing in winter: sources, variations
817 and processes, Atmos. Chem. Phys., 16, 8309-8329, 10.5194/acp-16-8309-2016, 2016.
- 818 Sun, Y. L., Zhang, Q., Anastasio, C., and Sun, J.: Insights into secondary organic aerosol formed via
819 aqueous-phase reactions of phenolic compounds based on high resolution mass spectrometry, Atmos.
820 Chem. Phys., 10, 4809-4822, 10.5194/acp-10-4809-2010, 2010.
- 821 Sun, Y. L., Wang, Z. F., Fu, P. Q., Yang, T., Jiang, Q., Dong, H. B., Li, J., and Jia, J. J.: Aerosol
822 composition, sources and processes during wintertime in Beijing, China, Atmos. Chem. Phys., 13,
823 4577-4592, 10.5194/acp-13-4577-2013, 2013b.
- 824 Tie, X., Huang, R.-J., Cao, J., Zhang, Q., Cheng, Y., Su, H., Chang, D., Pöschl, U., Hoffmann, T.,
825 Dusek, U., Li, G., Worsnop, D. R., and O'Dowd, C. D.: Severe pollution in China amplified by
826 atmospheric moisture, Sci. Rep., 7, 15760, 10.1038/s41598-017-15909-1, 2017.
- 827 Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.: Interpretation of
828 organic components from Positive Matrix Factorization of aerosol mass spectrometric data, Atmos.
829 Chem. Phys., 9, 2891-2918, 10.5194/acp-9-2891-2009, 2009.
- 830 Wang, G., Zhang, R., Gomez, M. E., Yang, L., Levy Zamora, M., Hu, M., Lin, Y., Peng, J., Guo, S.,
831 Meng, J., Li, J., Cheng, C., Hu, T., Ren, Y., Wang, Y., Gao, J., Cao, J., An, Z., Zhou, W., Li, G., Wang,



- 832 J., Tian, P., Marrero-Ortiz, W., Secrest, J., Du, Z., Zheng, J., Shang, D., Zeng, L., Shao, M., Wang, W.,
833 Huang, Y., Wang, Y., Zhu, Y., Li, Y., Hu, J., Pan, B., Cai, L., Cheng, Y., Ji, Y., Zhang, F., Rosenfeld, D.,
834 Liss, P. S., Duce, R. A., Kolb, C. E., and Molina, M. J.: Persistent sulfate formation from London Fog
835 to Chinese haze, *P. Natl. Acad. Sci. USA*, 113, 13630-13635, 10.1073/pnas.1616540113, 2016a.
- 836 Wang, J., Ge, X., Chen, Y., Shen, Y., Zhang, Q., Sun, Y., Xu, J., Ge, S., Yu, H., and Chen, M.: Highly
837 time-resolved urban aerosol characteristics during springtime in Yangtze River Delta, China: insights
838 from soot particle aerosol mass spectrometry, *Atmos. Chem. Phys.*, 16, 9109-9127,
839 10.5194/acp-16-9109-2016, 2016b.
- 840 Wang, J., Onasch, T. B., Ge, X., Collier, S., Zhang, Q., Sun, Y., Yu, H., Chen, M., Prévôt, A. S. H., and
841 Worsnop, D. R.: Observation of fullerene soot in eastern China, *Environ. Sci. Technol. Lett.*, 3,
842 121-126, 10.1021/acs.estlett.6b00044, 2016c.
- 843 Wang, X., Wang, W., Yang, L., Gao, X., Nie, W., Yu, Y., Xu, P., Zhou, Y., and Wang, Z.: The secondary
844 formation of inorganic aerosols in the droplet mode through heterogeneous aqueous reactions under
845 haze conditions, *Atmos. Environ.*, 63, 68-76, 10.1016/j.atmosenv.2012.09.029, 2012.
- 846 Wexler, A. S., and Johnston, M. V.: What have we learned from highly time-resolved measurements
847 during EPA's Supersites program, and related studies?, *J. Air Waste Manage.*, 58, 303-319,
848 10.3155/1047-3289.58.2.303, 2008.
- 849 Wu, D., Zhang, F., Ge, X., Yang, M., Xia, J., Liu, G., and Li, F.: Chemical and light extinction
850 characteristics of atmospheric aerosols in suburban Nanjing, China, *Atmosphere*, 8, 149,
851 10.3390/atmos8080149, 2017.
- 852 Xie, Y., Ding, A., Nie, W., Mao, H., Qi, X., Huang, X., Xu, Z., Kerminen, V.-M., Petäjä, T., Chi, X.,
853 Virkkula, A., Boy, M., Xue, L., Guo, J., Sun, J., Yang, X., Kulmala, M., and Fu, C.: Enhanced sulfate
854 formation by nitrogen dioxide: Implications from in situ observations at the SORPES station, *J.*
855 *Geophys. Res.*, 120, 2015JD023607, 10.1002/2015JD023607, 2015.
- 856 Xu, W., Han, T., Du, W., Wang, Q., Chen, C., Zhao, J., Zhang, Y., Li, J., Fu, P., Wang, Z., Worsnop, D.
857 R., and Sun, Y.: Effects of aqueous-phase and photochemical processing on secondary organic aerosol
858 formation and evolution in Beijing, China, *Environ. Sci. Technol.*, 51, 762-770,
859 10.1021/acs.est.6b04498, 2017a.
- 860 Xu, W., Sun, Y., Wang, Q., Du, W., Zhao, J., Ge, X., Han, T., Zhang, Y., Zhou, W., Li, J., Fu, P., Wang,
861 Z., and Worsnop, D. R.: Seasonal characterization of organic nitrogen in atmospheric aerosols using
862 high resolution aerosol mass spectrometry in Beijing, China, *ACS Earth Space Chem.*, 1, 673-682,
863 10.1021/acsearthspacechem.7b00106, 2017b.
- 864 Yan, C., Zheng, M., Sullivan, A. P., Shen, G., Chen, Y., Wang, S., Zhao, B., Cai, S., Desyaterik, Y., Li,
865 X., Zhou, T., Gustafsson, Ö., and Collett, J. L.: Residential coal combustion as a source of
866 levoglucosan in China, *Environ. Sci. Technol.*, 10.1021/acs.est.7b05858, 2018.
- 867 Yang, T., Sun, Y., Zhang, W., Wang, Z., Liu, X., Fu, P., and Wang, X.: Evolutionary processes and
868 sources of high-nitrate haze episodes over Beijing, Spring, *J. Environ. Sci.*, 54, 142-151,



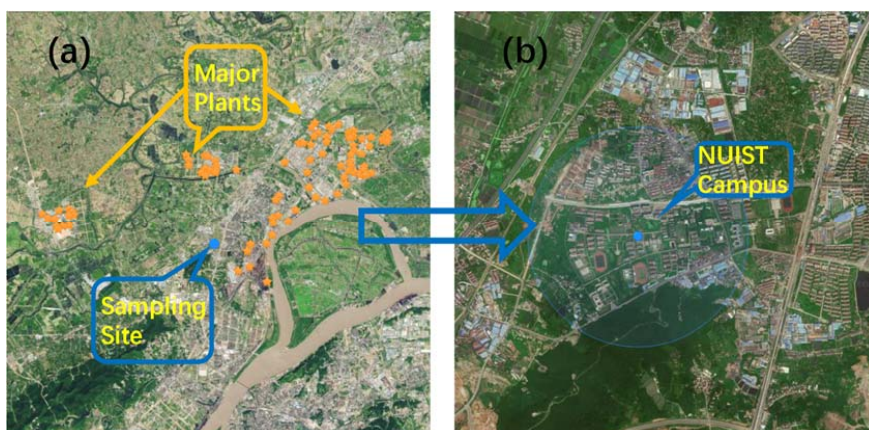
- 869 10.1016/j.jes.2016.04.024, 2017.
- 870 Ye, Z., Li, Q., Liu, J., Luo, S., Zhou, Q., Bi, C., Ma, S., Chen, Y., Chen, H., Li, L., and Ge, X.:
871 Investigation of submicron aerosol characteristics in Changzhou, China: Composition, source, and
872 comparison with co-collected PM_{2.5}, *Chemosphere*, 183, 176-185,
873 <https://doi.org/10.1016/j.chemosphere.2017.05.094>, 2017a.
- 874 Ye, Z., Li, Q., Ma, S., Zhou, Q., Gu, Y., Su, Y., Chen, Y., Chen, H., Wang, J., and Ge, X.: Summertime
875 day-night differences of PM_{2.5} components (inorganic ions, OC, EC, WSOC, WSON, HULIS, and
876 PAHs) in Changzhou, China, *Atmosphere*, 8, 189, 10.3390/atmos8100189, 2017b.
- 877 Ye, Z., Liu, J., Gu, A., Feng, F., Liu, Y., Bi, C., Xu, J., Li, L., Chen, H., Chen, Y., Dai, L., Zhou, Q., and
878 Ge, X.: Chemical characterization of fine particulate matter in Changzhou, China and source
879 apportionment with offline aerosol mass spectrometry, *Atmos. Chem. Phys.*, 2017, 2573-2592,
880 10.5194/acp-2016-883, 2017c.
- 881 Yu, L., Smith, J., Laskin, A., Anastasio, C., Laskin, J., and Zhang, Q.: Chemical characterization of
882 SOA formed from aqueous-phase reactions of phenols with the triplet excited state of carbonyl and
883 hydroxyl radical, *Atmos. Chem. Phys.*, 14, 13801-13816, 10.5194/acp-14-13801-2014, 2014.
- 884 Yu, L., Smith, J., Laskin, A., George, K. M., Anastasio, C., Laskin, J., Dillner, A. M., and Zhang, Q.:
885 Molecular transformations of phenolic SOA during photochemical aging in the aqueous phase:
886 competition among oligomerization, functionalization, and fragmentation, *Atmos. Chem. Phys.*, 16,
887 4511-4527, 10.5194/acp-16-4511-2016, 2016.
- 888 Zhang, C., Lu, X., Zhai, J., Chen, H., Yang, X., Zhang, Q., Zhao, Q., Fu, Q., Sha, F., and Jin, J.:
889 Insights into the formation of secondary organic carbon in the summertime in urban Shanghai, *J.*
890 *Environ. Sci.*, 10.1016/j.jes.2017.12.018, 2018a.
- 891 Zhang, H., Chen, S., Zhong, J., Zhang, S., Zhang, Y., Zhang, X., Li, Z., and Zeng, X.: Formation of
892 aqueous-phase sulfate during the haze period in China: Kinetics and atmospheric implications, *Atmos.*
893 *Environ.*, 10.1016/j.atmosenv.2018.01.017, 2018b.
- 894 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I., Alfarra, M. R., Takami,
895 A., Middlebrook, A. M., Sun, Y. L., Dzepina, K., Dunlea, E., Docherty, K., DeCarlo, P. F., Salcedo, D.,
896 Onasch, T., Jayne, J. T., Miyoshi, T., Shimojo, A., Hatakeyama, S., Takegawa, N., Kondo, Y.,
897 Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demetjian, K., Williams, P., Bower, K.,
898 Bahreini, R., Cottrell, L., Griffin, R. J., Rautiainen, J., Sun, J. Y., Zhang, Y. M., and Worsnop, D. R.:
899 Ubiquity and dominance of oxygenated species in organic aerosols in anthropogenically-influenced
900 Northern Hemisphere midlatitudes, *Geophys. Res. Lett.*, 34, L13801, 10.1029/2007gl029979, 2007.
- 901 Zhang, Q., Jimenez, J., Canagaratna, M., Ulbrich, I., Ng, N., Worsnop, D., and Sun, Y.: Understanding
902 atmospheric organic aerosols via factor analysis of aerosol mass spectrometry: a review, *Anal. Bioanal.*
903 *Chem.*, 401, 3045-3067, 10.1007/s00216-011-5355-y, 2011.
- 904 Zhang, R., Sun, X., Huang, Y., Shi, A., Yan, J., Nie, T., Yan, X., and Li, X.: Secondary inorganic
905 aerosols formation during haze episodes at an urban site in Beijing, China, *Atmos. Environ.*,



- 906 10.1016/j.atmosenv.2017.12.031, 2018c.
- 907 Zhang, Y., Tang, L., Croteau, P. L., Favez, O., Sun, Y., Canagaratna, M. R., Wang, Z., Couvidat, F.,
908 Albinet, A., Zhang, H., Sciare, J., Prévôt, A. S. H., Jayne, J. T., and Worsnop, D. R.: Field
909 characterization of the PM_{2.5} Aerosol Chemical Speciation Monitor: insights into the composition,
910 sources, and processes of fine particles in eastern China, *Atmos. Chem. Phys.*, 17, 14501-14517,
911 10.5194/acp-17-14501-2017, 2017.
- 912 Zhang, Y. J., Tang, L. L., Wang, Z., Yu, H. X., Sun, Y. L., Liu, D., Qin, W., Canonaco, F., Prévôt, A. S.
913 H., Zhang, H. L., and Zhou, H. C.: Insights into characteristics, sources, and evolution of submicron
914 aerosols during harvest seasons in the Yangtze River delta region, China, *Atmos. Chem. Phys.*, 15,
915 1331-1349, 10.5194/acp-15-1331-2015, 2015a.
- 916 Zhang, Y. J., Tang, L., Yu, H., Wang, Z., Sun, Y., Qin, W., Chen, W., Chen, C., Ding, A., Wu, J., Ge, S.,
917 Chen, C., and Zhou, H.-c.: Chemical composition, sources and evolution processes of aerosol at an
918 urban site in Yangtze River Delta, China during wintertime, *Atmos. Environ.*, 123, 339-349,
919 10.1016/j.atmosenv.2015.08.017, 2016.
- 920 Zhang, Y. W., Zhang, X. Y., Zhang, Y. M., Shen, X. J., Sun, J. Y., Ma, Q. L., Yu, X. M., Zhu, J. L.,
921 Zhang, L., and Che, H. C.: Significant concentration changes of chemical components of PM₁ in the
922 Yangtze River Delta area of China and the implications for the formation mechanism of heavy
923 haze-fog pollution, *Sci. Total Environ.*, 538, 7-15, 10.1016/j.scitotenv.2015.06.104, 2015b.
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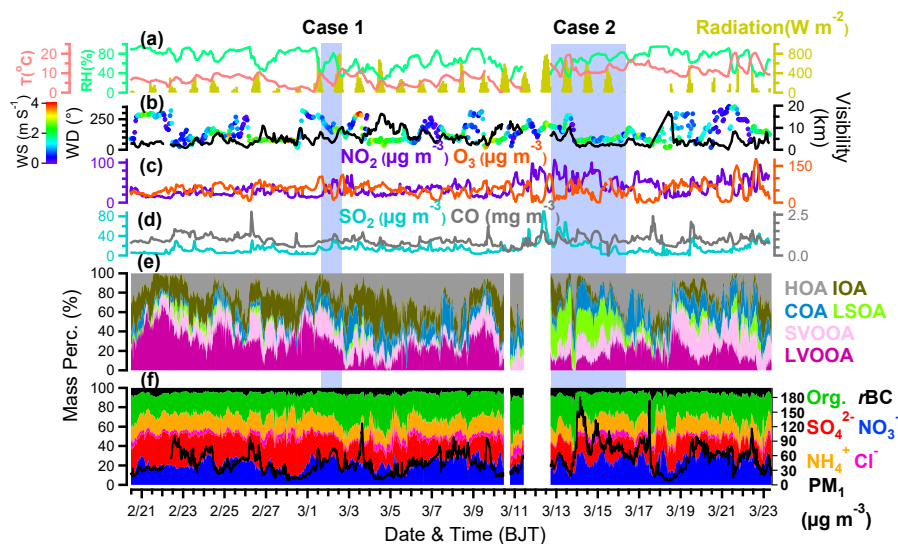
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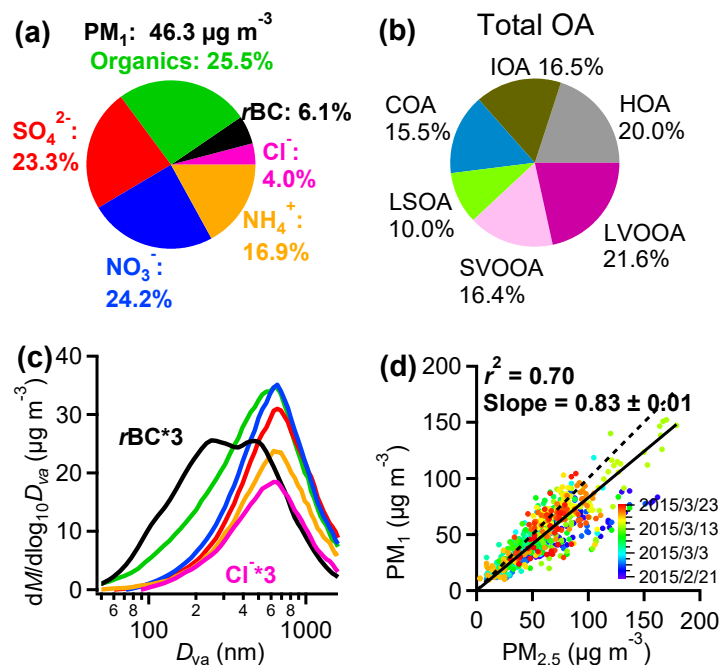
928 Figure 1. (a) The sampling site and its surroundings. The solid blue point is the
929 sampling site inside the campus of Nanjing University of Information Science and
930 Technology (NUIST) (b). The orange stars mark the positions of the major plants
931 adjacent to the site.

932



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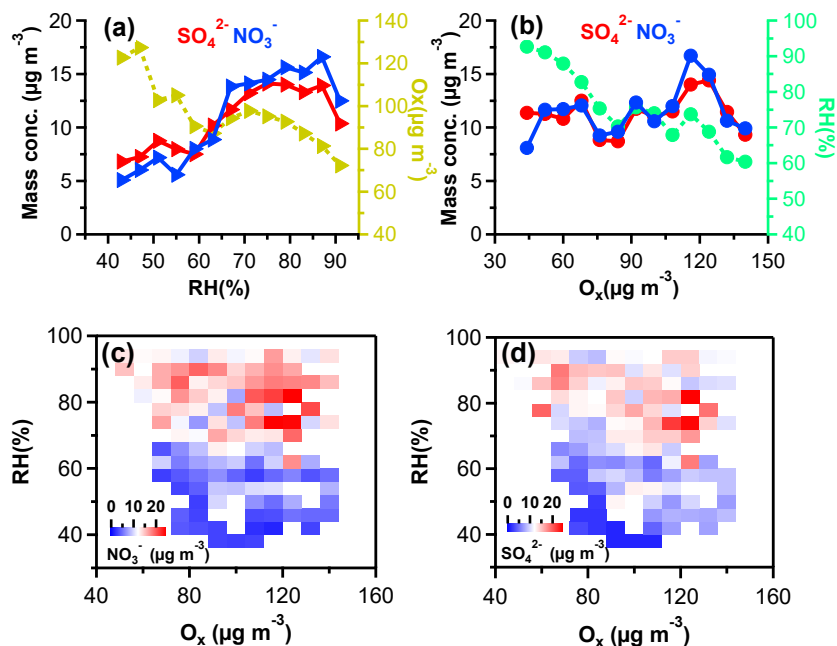
934 Figure 2. Time series of (a) air temperature (T), relative humidity (RH), and solar
935 radiation; (b) wind direction (WD) colored by wind speed (WS), and visibility (km);
936 (c-d) mass concentrations of NO_2 , O_3 , SO_2 and CO (hourly data); (e) mass
937 contributions (%) of the OA subcomponents including the hydrocarbon-like organic
938 aerosol (HOA), industry-related OA (IOA), cooking OA (COA), local secondary OA
939 (LSOA), semi-volatile oxygenated OA (SVOOA) and low-volatility oxygenated OA
940 (LVOOA); and (f) mass contributions (%) of refractory black carbon (rBC), total
941 organics, ammonium, chloride, sulfate and nitrate. Two specific episodes are marked
942 in light purple. (BJT, Beijing Time).



943

944 Figure 3. (a) Campaign-averaged mass fractions of rBC, organics, sulfate, nitrate,
945 ammonium and chloride to the total PM₁; (b) campaign-averaged mass fractions of
946 the six OA factors to the total OA; (c) mass-based campaign-average size distributions
947 of rBC ($\times 3$), organics, sulfate, nitrate, ammonium and chloride ($\times 3$); (d) scatter plots
948 of the PM₁ determined in this study vs. the PM_{2.5} from government monitoring station
949 (colored by time).

950



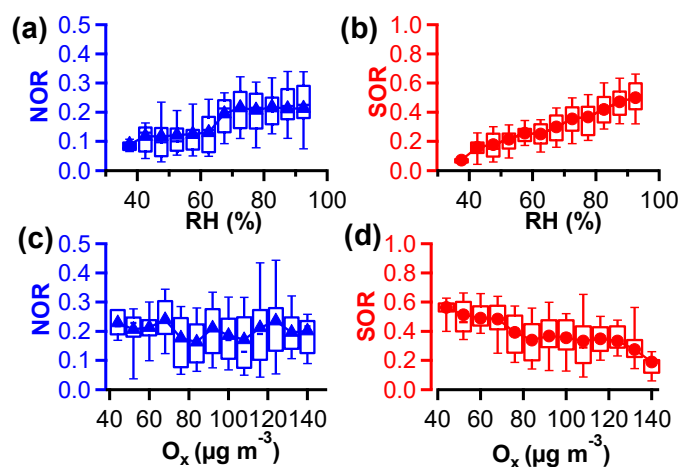
951

952 Figure 4. (a) Sulfate, nitrate and O_x concentrations vs. RH (5% increment); (b) sulfate,
953 nitrate concentrations and RH vs. O_x concentrations ($8 \mu\text{g}/\text{m}^3$ increment); (c-d) RH- O_x
954 image plots colored by sulfate and nitrate concentrations, respectively.

955



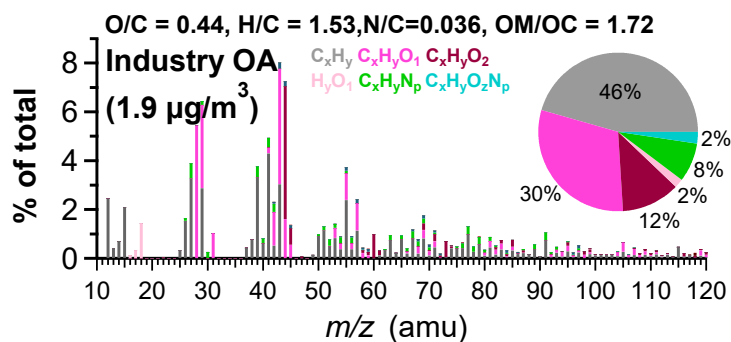
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957

958 Figure 5. (a-b) Variations of nitrogen oxidation ratio (NOR)(a) and sulfur oxidation
 959 ratio (SOR) (b) as a function of RH (5% increment); and (c-d) NOR and SOR
 960 variations as a function of O_x concentrations (8 μg/m³ increment) (the lines and solid
 961 triangles are the mean values, the lines in the boxes are the median values, the upper
 962 and lower boundaries of the boxes indicate the 75th and 25th percentiles, and the
 963 whiskers above and below the boxes indicate the 90th and 10th percentiles).

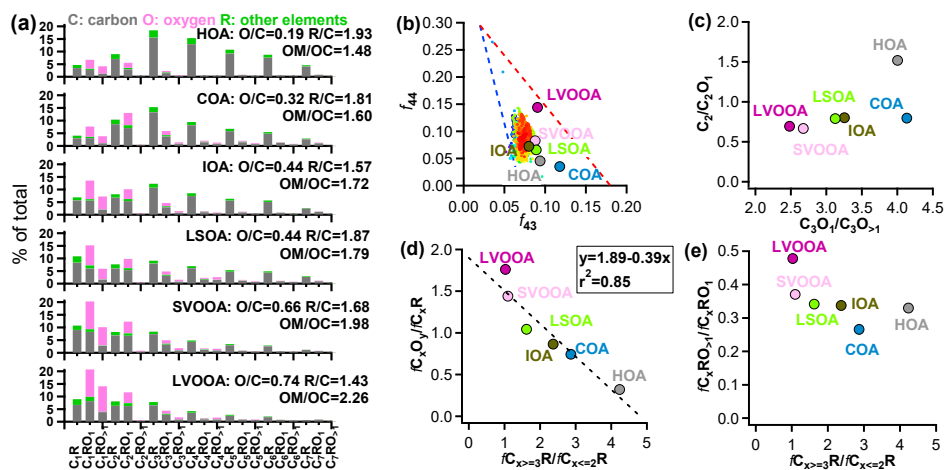
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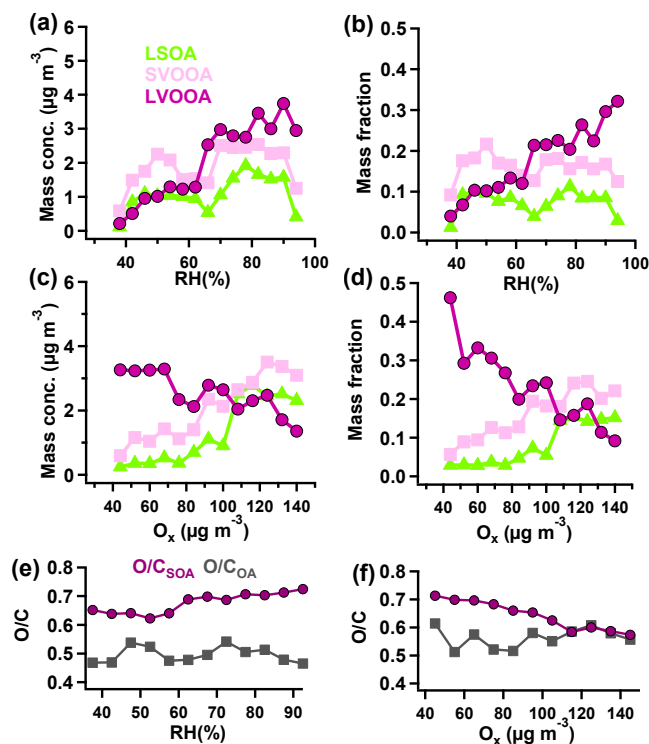
966 Figure 6. High-resolution mass spectrum of the industry-related OA (IOA) colored by
 967 six ion categories (the inset pie shows the relative mass contributions of the six ion
 968 categories).

969



970

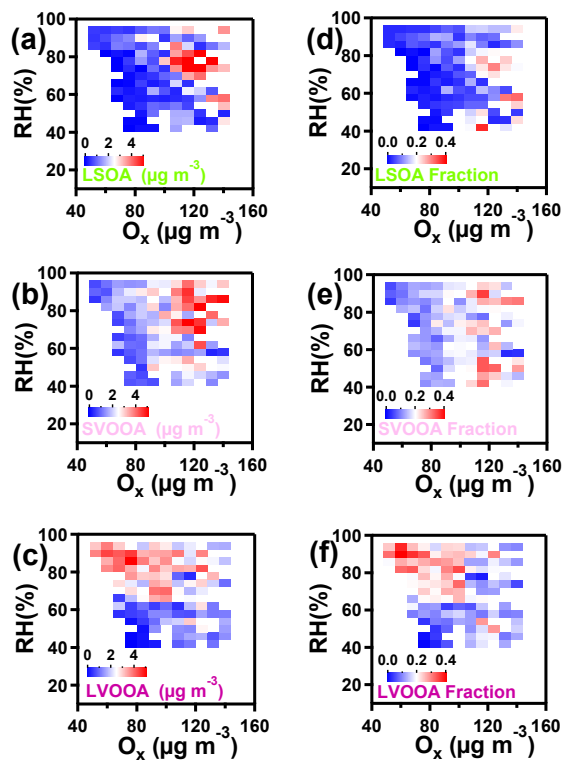
971 Figure 7. (a) Mass spectra of the six OA factors resorted by the ion groups with
 972 different carbon numbers (1-7) and oxygen numbers (0, 1, >1); (b) triangle plot of f_{44}
 973 vs. f_{43} for all OA (colored by time) and the six OA factors; and distributions of the six
 974 OA factors in the C_2/C_2O_1 vs. $C_3O_1/C_3O_{>1}$ space (c), the $f_{C_xO_y}/f_{C_xR}$ vs. $f_{C_x \geq 3R}/f_{C_x \leq 2R}$
 975 space (d), and the $f_{C_xRO_{>1}}/f_{C_xRO_1}$ vs. $f_{C_x \geq 3R}/f_{C_x \leq 2R}$ space (Meanings of the ion
 976 groups are described in the main text).



977

978 Figure 8. Mass concentrations (a) and fractional contributions (b) of the LSOA,
979 SVOOA and LVOOA as a function of RH (5% increment); mass concentrations (c)
980 and fractional contributions of LSOA, SVOOA and LVOOA as a function of O_x
981 concentrations ($8 \mu\text{g}/\text{m}^3$ increment); and the O/C values of total OA ($\text{O}/\text{C}_{\text{OA}}$) and SOA
982 ($\text{O}/\text{C}_{\text{SOA}}$) against RH (e) and O_x concentrations (f).

983



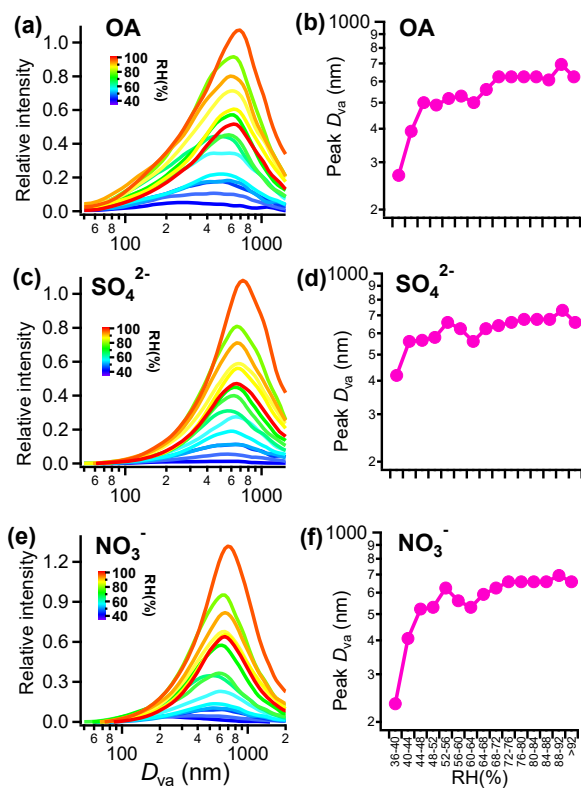
984

985 Figure 9. RH-O_x image plots colored by the mass concentrations of LSOA, SVOOA

986 and LVOOA (a-c), and the mass fractions of LSOA, SVOOA and LVOOA to total OA

987 (d-f).

988



989

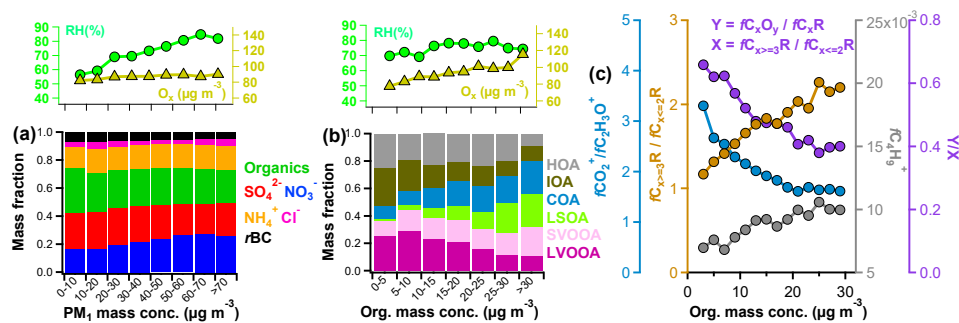
990 Figure 10. Average size distributions and the peak D_{va} of OA (a-b), sulfate (c-d) and

991 nitrate (e-f) at different RH bins (4% increment).

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993



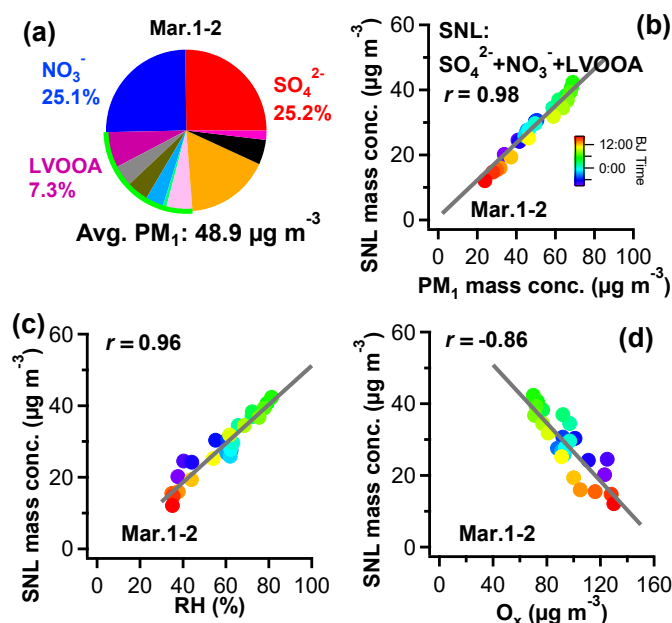
994

995 Figure 11. (a) Mass fractions of the six PM₁ components, and corresponding average
 996 RH values and O_x concentrations at different PM₁ concentrations bins (10 µg/m³
 997 increment); (b) mass fractions of the six OA factors, and corresponding average RH
 998 values and O_x concentrations at different OA concentration bins (5 µg/m³ increment);
 999 (c) $f\text{CO}_2^+/f\text{C}_2\text{H}_3\text{O}^+$, $f\text{C}_{x\geq 3}\text{R}/f\text{C}_{x\leq 2}\text{R}$, $f\text{C}_4\text{H}_9^+$ and $(f\text{C}_x\text{O}_y/f\text{C}_x\text{R})/(f\text{C}_{x\geq 3}\text{R}/f\text{C}_{x\leq 2}\text{R})$ as a
 1000 function of OA concentrations.

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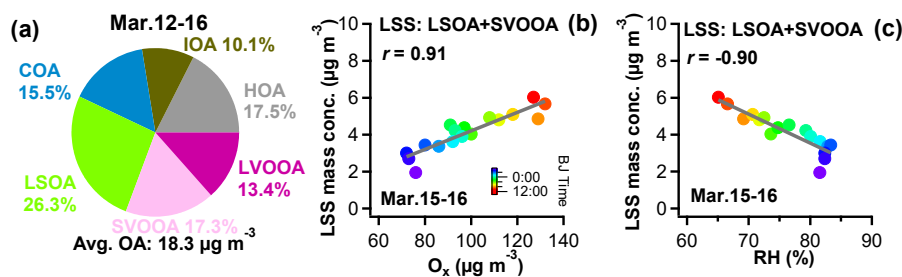
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1003

1004 Figure 12. Case 1 (3:30 pm March 1 – 3:30 pm March 2): (a) average mass
 1005 percentages of different species; scatter plots of SNL concentrations (the sum of
 1006 sulfate, nitrate and LVOOA concentrations) vs. PM₁ concentrations (b), RH (c) and O_x
 1007 concentrations (d) (colored by time).

1008



1009

1010 Figure 13. Case 2 (6:30 pm March 12 – 8:00 am March 16): (a) average mass
 1011 contributions of the six OA factors to OA; scatter plots of LSS concentrations (the
 1012 sum of LSOA and SVOOA concentrations) vs. O_x concentrations (b) and RH (c)
 1013 (colored by time) (the data was only for 12:00 pm March 15 – 8:00 am March 16).

1014