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2 **Diagnosis of dust- and haze pollution-impacted PM₁₀, PM_{2.5},**
3 **and PM₁ aerosols observed at Gosan Climate Observatory**

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22



23 **Abstract**

24

25 In East Asia, soil dust is a major component of aerosols and is mixed with various pollutants
26 during transport, resulting in large uncertainty in climate and environmental impact
27 assessment and relevant policymaking. To diagnose the influence of soil dust and
28 anthropogenic pollution on bulk aerosol, we conducted long-term measurements of mass,
29 water-soluble ions, and carbonaceous compounds of PM_{10} , $PM_{2.5}$, and PM_1 at Gosan Climate
30 Observatory, South Korea, from August 2007 to February 2012. The principle component
31 analyses of all measured species reveal that the impact of anthropogenic pollution, soil dust,
32 and agricultural fertilizer accounts for 46%, 16%, and 9% of the total variance, respectively.
33 Particularly, the loadings of agricultural component were high in the warmer months with the
34 least occurrence of high concentration events and have increased over time. In mode analysis
35 of PM_{10} , $PM_{2.5}$, and PM_1 mass concentrations, the mean + σ was comparable to the 90th
36 percentile and thus, suggested as a robust criterion that determines the substantial impact of
37 soil dust and haze pollution on particulate matter. The results of this study imply that non-
38 combustion sources such as soil dust will impose constraints to the reduction of $PM_{2.5}$ as well
39 as PM_{10} concentrations. In addition, questions are raised as to whether the yearly average
40 concentration is suitable for environmental standard in northeast Asian region.

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43 Introduction

44

45 Dust particles are dominant atmospheric aerosols and account for more than 60% of the total
46 global dry aerosol mass burden (Textor et al., 2006). The emission of dust ranges from 1000
47 to 3000 Tg yr⁻¹ (Zender et al., 2004). Dust particles are abundant in coarse mode, which is
48 represented by PM₁₀. Recently, a new type of dust particle has been observed in submicron
49 size that was long-range transported from dry lake deposits in northern China. These particles
50 are more abundant in salts and mineral constituents compared with typical soil dust (Shang et
51 al., 2018). The high-mass loading of dust adversely affects air quality and causes climate
52 change. In regional to global earth's environment, mineral dust contributes to a radiative
53 forcing of $-0.3 \sim +0.1 \text{ W m}^{-2}$ with a large uncertainty (IPCC, 2007, 2013) because emission
54 source, chemical and mineralogical composition, and particle size vary in a wide range
55 (Choobari et al., 2014). Dust particles have also been predicted to reduce the rate of ozone
56 production (Dickerson et al., 1997) and promote new particle formation and growth by
57 mixing with other pollutants (Nie et al., 2014; He et al., 2014). Their role in climate change is
58 further highlighted by modifying cloud microphysical process via cloud droplet activation
59 (Bègue et al., 2015) and CO₂ uptake via ocean fertilization (Pabortsava, et al., 2017).

60

61 A large amount of dust is found over arid regions in North Africa, North China and South
62 Mongolia (Choobari et al., 2014). In East Asian desert areas, soil type significantly varies
63 from arid and semi-arid deserts to loess deposits and dry lake deposits, depending on the
64 formation process and climate (Zhang et al., 2009). Dust particles generated from these
65 regions are different in chemical and mineralogical composition (Cheng et al., 2012; Kunwar
66 and Kawamura, 2014; Liu et al., 2011; Tao et al., 2014; Su and Toon, 2011), leading to
67 different environmental effects. For example, African dust contains more iron oxides mineral
68 in the form of hematite as compared with those from Asia (Formenti et al., 2011; 2014),
69 implying greater light absorption (Zhang et al., 2015). Among atmospheric aerosols generated
70 from variety of emissions (Geng et al., 2014). EC is a main species absorbing light and
71 measured as black carbon (BC) (Han et al., 2010; Saleh et al., 2014). In Northeast Asia,
72 aerosols from various sources are often mixed together while being transported (Kim et al.,



73 2007) and dust plumes have been identified as enhanced concentrations of sulfate, nitrate,
74 ammonium, OC or EC over the Korean peninsula (Shin et al., 2015).

75

76 An array of AERONET and satellite observations is a vital tool for investigating atmospheric
77 aerosol at regional and global scales (Zhuang et al., 1992; Zhang et al., 2003; Jin et al., 2016;
78 Nazari et al., 2016). Because these techniques measure the extinction of bulk aerosols, it is
79 crucial to accurately estimate the optical property of major aerosol constituents (e.g., Kim et
80 al., 2007) in order to assess the anthropogenic contribution to total aerosol burden and its
81 climate effect using remote sensing data. However, it is still a big challenge to estimate the
82 optical property of main aerosol types especially dust particles in East Asia, because their
83 property is not only dependent on source regions (Huang et al., 2014) but also modified
84 during transport through aging and mixing with pollutants (McFarlane et al., 1992; Von
85 Salzen et al., 2005; Bäumer et al., 2007; Kim et al., 2011). For instance, the mixture of dust
86 and pollutant caused a significant increase of radiative forcing by 0.06 Wm^{-2} (Zhang et al.,
87 2013a; Mishra et al., 2010; Li et al., 2012). In this context, it is important to figure out the
88 extent of dust particles in the atmosphere and distinguish them from bulk aerosol particles.
89 For instance, the ratio of $\text{PM}_{10}/\text{PM}_{2.5}$ has been employed to eliminate the effect of soil dust
90 from $\text{PM}_{2.5}$ when determining the mass absorption coefficient of OC (Chung et al., 2012).

91

92 In this study, to diagnose the impact of airborne dust and anthropogenic pollution on
93 atmospheric particulate matter and find out feasible criteria, the five-year measurements of
94 mass, water-soluble ions, and carbonaceous compounds for PM_{10} , $\text{PM}_{2.5}$, and PM_1 were
95 analyzed using statistical methods. The long-term and composite measurements of particulate
96 matter in different sizes are scarce in the study region. The results of mass mode analysis and
97 principle component analysis will provide insight into the role and significance of mineral
98 dust and haze pollution on particulate matter and further its control strategies in northeast
99 Asia.

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102 **1 Methodology**

103

104 Aerosol samples were collected separately for PM₁, PM_{2.5}, and PM₁₀ onto 37 mm Teflon and
105 Quartz filters (Pall, Corp.) using sharp-cut cyclones (URG, USA) at the Gosan Climate
106 Observatory (GCO) from 2007 to 2012. Sampling was undertaken for a period of 24 h from
107 10:00 to 10:00 the next day. A total of 152 sets of samples were collected and analyzed for
108 water-soluble inorganic ions and carbonaceous compounds. Details about the measurement
109 methodology can be found in Lim et al. (2012, 2014). During the five-year period, five dust
110 and eleven haze events were recorded across Korea by the Korea Meteorological
111 Administration (KMA) mostly during the cold seasons from late fall to spring (Fig. 1).

112

113 Teflon filters were conditioned for 24 h in desiccators (SANPLATEC, Japan) under a relative
114 humidity of approximately 30%–40% and weighed before and after sampling using an
115 analytical balance (Denver, Germany). Water-soluble species were extracted from the filters
116 into a solution comprising a mixture of 19 mL distilled water and 1 mL methanol. Water-
117 soluble ions, including Cl⁻, NO₃⁻, SO₄²⁻, Na⁺, NH₄⁺, K⁺, Mg²⁺, and Ca²⁺ were analyzed via
118 ion chromatography (IC 25, Dionex, USA). For this analysis, 500 µL of sample was injected
119 by an auto-sampler into the AG11 and AS11 columns for anions or CG11 and CS11 columns
120 for cations (Dionex, USA). The eluent and suppressor were 39 mM KOH and
121 ASRSIIULTRA-4 mm and 20 mM MSA and CSRSIIULTRA-4 mm for anions and cations,
122 respectively. Finally, concentrations were determined using a conductivity detector (Dionex,
123 USA), which was calibrated against eight aqueous standards. The detection limit, defined as
124 3σ, was approximately 0.01–0.09 µg/m³.

125

126 Carbonaceous components were measured at the Desert Research Institute (the
127 thermal/optical reflectance (TOR) method, Reno, NV, USA) using the Interagency
128 Monitoring of Protected Visual Environments (IMPROVE) TOR protocol. OC comprising
129 OC1, OC2, OC3, and OC4 was determined at 120°C, 250°C, 450°C, and 550°C, respectively,
130 in a He atmosphere. EC was analyzed as EC1, EC2, and EC3 at 550°C, 700°C, and 850°C,
131 respectively, after introducing 2% O₂/ 98% He. Pyrolyzed OC (OP) was measured in the



132 O₂/He atmosphere before the reflected light returned to its initial value (Lim et al., 2012;
133 2014).

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136 **2 Measurement overview**

137

138 From August 2007 to December 2012, the average PM₁₀, PM_{2.5}, and PM₁ mass concentrations
139 of all measurements were 30 µg/m³, 19 µg/m³, and 14 µg/m³, respectively (Table 1). The
140 PM₁₀ mass was almost equally partitioned between <1 µm and 1 – 10 µm. Moreover, the
141 mass of particles between 1 µm and 2.5 µm was considerable and comprises 26% of PM_{2.5}
142 mass. As summarized in Table 1, SO₄²⁻ and OC were the most abundant, followed by NH₄⁺
143 and NO₃⁻. These four species accounted for 48%, 58%, and 69% of the PM₁₀, PM_{2.5}, and PM₁
144 mass, respectively. Of these species, SO₄²⁻, NH₄⁺, and EC were pre-dominant in PM₁, which
145 corresponds to more than 75% of those in PM₁₀. In comparison, about 65% of OC was
146 partitioned into PM₁. It was even less for NO₃⁻ as 33%. It is well known that NO₃⁻ is more
147 abundant in coarse mode particles due to high affinity to soil mineral. It is also noteworthy
148 that a substantial amount of OC and EC (20 %) was associated with particles between 1 µm
149 and 2.5 µm. Of OC sub-components, OC3 and OC4 were mainly associated with coarse
150 particles. It is evident that Na⁺ and Cl⁻ were highly enriched in coarse particles between 2.5
151 µm and 10 µm.

152

153 The high PM₁₀ mass was usually observed in the spring along with increased concentrations
154 of Ca²⁺ and Mg²⁺ (Fig. 1). In comparison, the mass of PM₁ was higher in the cold season (fall
155 to winter) when the concentrations of SO₄²⁻, NO₃⁻, NH₄⁺, K⁺, OC, and EC were highly
156 elevated. High PM_{2.5} concentrations were observed in both winter and spring periods. Overall,
157 the three particle masses and their major constituents were highly elevated when dust and
158 haze events occurred.

159



160 Five dust events took place in March, May, November, and December for the entire
161 experiment period. Upon dust incidence, the daily average mass concentrations of PM₁₀ and
162 PM_{2.5} were enhanced by 3 and 2 times, respectively. In comparison, concentrations of mass,
163 secondary ions, and carbonaceous compounds were elevated more than two times in PM₁
164 during the haze events from October to April. On March 20, 2010, dust and haze event
165 occurred concurrently, leading to a maximum PM₁₀ concentration of 199 µg/m³. PM₁₀
166 concentrations were occasionally elevated without an official report of KMA on dust
167 occurrence. For instance, in March 2008, the concentrations of PM₁₀ mass, Ca²⁺, and Mg²⁺
168 were higher than those of dust event days.

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171 **3 PCA analysis of PM₁₀, PM_{2.5}, and PM₁**

172

173 Principle component analysis (PCA) was conducted for all measured species of PM₁₀, PM_{2.5},
174 and PM₁ aerosols including water-soluble ions, OC, EC, and mass for the whole period. PCA
175 analysis identifies the correlation of variables through orthogonal transformation and
176 summarizes the main characteristics of measurement data set. In PCA analysis, two
177 components are usually selected. In this study, the principle component 1 and 2 accounted for
178 more than 60% of the total variance of PM₁₀, PM_{2.5}, and PM₁. The principle component 1
179 (PC1) was composed of high loadings for SO₄²⁻, NO₃⁻, NH₄⁺, K⁺, OC, and EC, especially in
180 PM₁ (Fig. 2). These six species contributed almost equally to the PC1, which explained 46%
181 of the total variance. In contrast, the principle component 2 (PC2) explained 16% of the total
182 variance and was characterized by high loadings for Na⁺, Cl⁻, Mg²⁺ and Ca²⁺, mainly in PM₁₀
183 and PM_{2.5}. Interestingly, the principle component 3 (PC3) comprising 9% of the total variance,
184 was associated with high loadings for NH₄⁺ and Ca²⁺, particularly in PM₁₀ (Fig. 2 and 3).
185 These three independent factors explain more than 70% of the total variance.

186

187 As the most abundant species, SO₄²⁻ and NO₃⁻ concentrations were highly correlated with
188 PC1 loadings for all three-size particles (Fig. 3), confirming that the PC1 represents the
189 influence of anthropogenic pollution sources (Zhang et al., 2013b). So were OC2 and EC1



190 that have been reported to originate from biomass combustion sources (e.g., Lim et al., 2012).
191 In PC2, the loadings for Ca^{2+} , Mg^{2+} , Na^+ , and Cl^- were the highest and well correlated with
192 OC4 concentration in $\text{PM}_{2.5}$ and PM_{10} , which used to be elevated upon dust events (Lim et al.,
193 2012). In saline dust, the concentrations of Ca^{2+} , Na^+ , and Cl^- were enhanced concurrently
194 with OC sub-component (Zhang et al., 2014; Shang et al., 2018; O'Dowd et al., 2004; Griffith
195 et al., 2010). The sea-salt contribution of Ca^{2+} was estimated to be 12% in $\text{PM}_{2.5}$ and 19% in
196 PM_{10} , assuming that sodium was derived solely from sea salt. In this study, the measurements
197 of water-soluble ions demonstrate that the contribution of sea salt species was found to reach
198 the maximum in summer when aerosol loading is at its minimum under influence of marine
199 air. Thus, the PC2 represents the impact of dust particles including alkaline soils. NH_4^+
200 concentration was moderately related to PC3 loadings in $\text{PM}_{2.5}$ and PM_{10} . In particular, a
201 relatively good correlation of NH_4^+ with Ca^{2+} in PM_{10} indicates the agricultural influence due
202 to fertilizer use. It is noteworthy that PC3 loading was high in spring and summer when the
203 concentrations of particulate matter were low with reduced continental outflows. In addition,
204 the PC3 loadings increased with time, reaching to the highest in 2010. The recent studies also
205 reported that in China, NH_3 emission was increased due to fertilizer application and NH_4^+
206 concentration was higher in spring and summer than the other seasons (Warner et al, 2017;
207 Kang et al., 2016).

208

209 Therefore, the three principle components manifest the main sources of particulate matters in
210 the study region. As anthropogenic sources, PC1 is predominant in PM_1 and $\text{PM}_{2.5}$. PC2
211 demonstrates the influence of soil dust on PM_{10} and $\text{PM}_{2.5}$. Fertilizer use is likely responsible
212 for the variance of PC3. In order to estimate the contribution of these three factors to the mass
213 of PM_{10} , $\text{PM}_{2.5}$, and PM_1 at GCO, multi-linear regression analysis was conducted using factor
214 loadings, of which result is given below:

215

$$216 \quad \text{PM}_{10} (\mu\text{g}/\text{m}^3) = 31.1 + 4.7 \text{PC1} + 3.7 \text{PC2} + 4.1 \text{PC3} \quad (r = 0.89, P = 0.03) \quad (1)$$

$$217 \quad \text{PM}_{2.5} (\mu\text{g}/\text{m}^3) = 19.2 + 3.1 \text{PC1} + 0.4 \text{PC2} + 1.9 \text{PC3} \quad (r = 0.95, P = 0.03) \quad (2)$$

$$218 \quad \text{PM}_1 (\mu\text{g}/\text{m}^3) = 14.8 + 2.5 \text{PC1} - 0.7 \text{PC2} + 1.6 \text{PC3} \quad (r = 0.93, P = 0.04) \quad (3)$$

219 , where PC1, PC2, and PC3 are factor loadings.



220

221 The intercepts of these three equations are equivalent to the average concentrations for PM₁₀,
222 PM_{2.5}, and PM₁ (Table 1 and 2). It confirms that the three PCs are sufficient enough to
223 explain the variation of aerosol masses observed at GCO. It is evident that PC1 is a dominant
224 factor determining the particulate mass of PM₁₀ (63%) as well as PM₁ (99%) and PM_{2.5} (90%).
225 PC2 was most evident in PM₁₀ (36%) and not negligible in PM_{2.5} (9%). It is worthy
226 emphasizing that NH₄⁺ factor was distinguished as PC3, even though its contribution was the
227 least. In addition, the very small or negative loading of PC2 for PM_{2.5} and PM₁ suggests the
228 scavenging of anthropogenic pollutants on dust particles.

229

230 **4 Diagnosis of dust and haze**

231

232 While soil dust has been recognized as a main driver for high PM₁₀ mass in northeast Asia
233 (Yang et al., 2009), air pollution events are typically distinguished by the concentrations of
234 PM_{2.5} (EPA, 2012). In Korea, the aerosol mass concentrations have been often elevated upon
235 Asian dust or haze occurrence. In this context, mode analysis of PM₁₀, PM_{2.5}, and PM₁ mass
236 concentrations was conducted to diagnose the impact of dust and haze particles on particulate
237 matter. The frequency distributions of all PM₁₀, PM_{2.5}, and PM₁ measurements are shown in
238 Figure 4. For the three-size aerosol masses, the main-mode concentrations are comparable to
239 the median concentrations (Table 3 and Fig. 4). The main-mode concentration of PM₁₀ and
240 PM_{2.5} was 25 µg/m³ and 16 µg/m³, respectively, which are much lower than those of national
241 standard of annual mean of 50 µg/m³ and 25 µg/m³, respectively. The main-mode
242 concentration of PM₁ (11 µg/m³) was similar to the air quality guideline of the World Health
243 Organization (WHO) for PM_{2.5} (10 µg/m³) (WHO, 2006). Of PM_{2.5} mass, the contribution of
244 mineral dust was estimated to be ~10% in previous section, which is equivalent to about 2
245 µg/m³.

246

247 The mean concentrations of the three types of particulate matters were higher than their
248 median and main-mode concentrations and the standard deviations were comparable to the
249 median concentrations. These results show that mass concentrations varied in a wide range



250 due to high concentration events. For PM_{10} , the mean+ σ of $52 \mu\text{g}/\text{m}^3$ was close to the national
251 standard of PM_{10} annual average concentration. While the mean+ σ concentration of $PM_{2.5}$
252 was higher by 28% than the national standard of $25 \mu\text{g}/\text{m}^3$, the mean+ σ of PM_1 ($25 \mu\text{g}/\text{m}^3$)
253 met the annual standard of $PM_{2.5}$ concentration. The mean+ σ concentrations of PM_{10} , $PM_{2.5}$,
254 and PM_1 were commensurate with the 90th percentiles that generally represent the highest
255 concentration of the long-term measurements.

256

257 In Korea, dust occurrence is determined by eye observation and haze is recorded when RH is
258 less than 75 % and visibility is between 1 km and 10 km. The concentrations of individual
259 dust and haze samples are presented in the bottom of Figure 4. While the PM_{10} and $PM_{2.5}$
260 concentrations of five dust events are placed in the range above the mean+ σ , all of the high
261 PM_{10} concentrations were not observed on dust days. In contrast, the concentrations of all
262 haze samples were over the mean+ σ of PM_1 (Fig. 4).

263

264 When PM_{10} and $PM_{2.5}$ mass belong to the top 10%, their Ca^{2+} and Mg^{2+} concentrations were
265 also within the highest 10 % of the entire measurements. Particularly, Ca^{2+} concentration (0.7
266 $\mu\text{g}/\text{m}^3$) was 3 times as high as the average concentration for both PM_{10} and $PM_{2.5}$ (Table 1),
267 implying that dust effect is not negligible in $PM_{2.5}$.

268

269 At GCO, the five-year measurements of aerosol mass and chemical composition reveal that
270 the top 10 % of PM_{10} , $PM_{2.5}$, and PM_1 mass was affected by dust or haze plumes and their
271 effect is traceable by the 90th percentile mass concentrations of particulate matter. If PM_{10} or
272 $PM_{2.5}$ mass concentrations are above the 90th percentile, airborne dust particles played a
273 substantial role in mass enhancement, regardless of the occurrence of event. Likewise,
274 anthropogenic pollution is a main driver for enhanced PM_1 and $PM_{2.5}$ mass concentrations if
275 their concentrations are greater than the 90th percentile. For $PM_{2.5}$, the impact of mineral dust
276 should be considered in northeast Asia region downwind of the dust belt.

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278

279 **5 Conclusions**

280

281 At GCO, filter samples for PM₁, PM_{2.5}, and PM₁₀ were collected and their mass, water-
282 soluble inorganic ions and carbonaceous compounds were analyzed from 2007 to 2012. For
283 the entire period, the average concentrations of PM₁₀, PM_{2.5}, and PM₁ were 30, 19, and 14
284 μg/m³, respectively. PM_{2.5} accounted for 63% of PM₁₀, while PM₁ comprised 74% of PM_{2.5}
285 on average.

286

287 From the principle component analysis using all measured species for PM₁₀, PM_{2.5}, and PM₁,
288 the three principle components (PC1, PC2, and PC3) were distinguished, which explained
289 46%, 16%, and 9% of the total variances, respectively. The PC1 representing the effect of
290 anthropogenic pollution was characterized by high loadings of SO₄²⁻, NO₃⁻, NH₄⁺, K⁺, OC,
291 and EC. The PC2 was distinct with high loadings for Ca²⁺ and Mg²⁺ that originate from soil
292 dust. Although the contribution was low, the PC3 was significant for two reasons. First, the
293 loadings of PC3 showed an increasing tendency over time. In addition, the PC3 loadings were
294 discernible during warm season, in contrast to other two components that explain the
295 variations of mass and major constituents of aerosol during cold season. The multiple
296 regression using the three PC loadings shows that the anthropogenic pollution accounted for
297 99 % and 63 % of PM₁ and PM₁₀ mass variation, respectively. The effect of soil dust was the
298 largest on PM₁₀ (36%) and not negligible on PM_{2.5} (~10%).

299

300 The mode analysis of PM₁₀, PM_{2.5}, and PM₁ mass concentrations demonstrates that the main
301 mode was commensurate with the median concentration and the mean + σ was comparable to
302 the concentration of the 90th percentile. It indicates that the average mass concentration is
303 highly susceptible to high-concentration episodes. Consequently, the mean+σ is suggested as
304 a robust criterion that determines the substantial impact of soil dust or pollution plumes on
305 PM₁₀, PM_{2.5}, and PM₁. Furthermore, the results of this study reveal that in northeast Asia,
306 non-combustion sources such as soil dust with impose constraints to the reduction of PM_{2.5} as
307 well as PM₁₀ concentrations and raise questions about the efficacy of yearly average
308 concentrations as environmental standards.



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319 **References**

320 Bäumer, D., Lohmann, U., Lesins, G., Li, J., and Croft, B.: Parameterizing the optical
321 properties of carbonaceous aerosols in the Canadian Centre for Climate Modeling and
322 Analysis Atmospheric General Circulation Model with impacts on global radiation and
323 energy fluxes, *J. Geophys. Res. Atmos.*, 112(D10), <https://doi.org/10.1029/2006JD007319>,
324 2007.

325 Bègue, N., Tulet, P., Pelon, J., Aouizerats, B., Berger, A., and Schwarzenboeck, A.: Aerosol
326 processing and CCN formation of an intense Saharan dust plume during the EUCAARI 2008
327 campaign, *Atmos. Chem. Phys.*, 15, 3497–3516, <https://doi.org/10.5194/acp-15-3497-2015>,
328 2015.

329 Cheng, M. C., You, C. F., Cao, J. J., and Jin, Z. D.: Spatial and seasonal variability of water-
330 soluble ions in PM_{2.5} aerosols in 14 major cities in China, *Atmos. Environ.*, 60, 182–192,
331 <https://doi.org/10.1016/j.atmosenv.2012.06.037>, 2012.

332 Choobari, O. A., Zawar-Reza, P., and Sturman, A.: The global distribution of mineral dust
333 and its impacts on the climate system: a review, *Atmos. Res.*, 138, 152–165,
334 <https://doi.org/10.1016/j.atmosres.2013.11.007>, 2014.

335 Chung, C. E., Kim, S. W., Lee, M., Yoon, S. C., and Lee, S.: Carbonaceous aerosol AAE
336 inferred from in-situ aerosol measurements at the Gosan ABC super site, and the implications
337 for brown carbon aerosol, *Atmos. Chem. Phys.*, 12, 6173–6184, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-12-6173-2012)
338 12-6173-2012, 2012.



- 339 Dickerson, R. R., Kondragunta, S., Stenchikov, G., Civerolo, K. L., Doddridge, B. G., and
340 Holben, B. N.: The impact of aerosols on solar ultraviolet radiation and photochemical smog,
341 *Science*, 278, 827–830, <https://doi.org/10.1126/science.278.5339.827>, 1997.
- 342 EPA: Our nation’s air: Status and trends through 2010, U.S. EPA Office of Air Quality
343 Planning and Standards, Research Triangle Park, NC, 2012.
- 344 Formenti, P., Schütz, L., Balkanski, Y., Desboeufs, K., Ebert, M., Kandler, K., Petzold, A.,
345 Scheuven, D., Weinbruch, S., and Zhang, D.: Recent progress in understanding physical and
346 chemical properties of African and Asian mineral dust, *Atmos. Chem. Phys.*, 11, 8231–8256,
347 <https://doi.org/10.5194/acp-11-8231-2011>, 2011.
- 348 Formenti, P., Caquineau, S., Desboeufs, K., Klaver, A., Chevaillier, S., Journet, E., and Rajot,
349 J. L.: Mapping the physico-chemical properties of mineral dust in western Africa:
350 mineralogical composition, *Atmos. Chem. Phys.*, 14, 10663–10686,
351 <https://doi.org/10.5194/acp-14-10663-2014>, 2014.
- 352 Geng, H., Hwang, H., Liu, X., Dong, S., and Ro, C. U.: Investigation of aged aerosols in size-
353 resolved Asian dust storm particles transported from Beijing, China, to Incheon, Korea, using
354 low-Z particle EPMA, *Atmos. Chem. Phys.*, 14, 3307–3323, <https://doi.org/10.5194/acp-14-3307-2014>, 2014.
- 356 Griffith, D. R., Martin, W. R., and Eglinton, T. I.: The radiocarbon age of organic carbon in
357 marine surface sediments, *Geochim. Cosmochim. Acta*, 74, 6788–6800,
358 <https://doi.org/10.1016/j.gca.2010.09.001>, 2010.
- 359 Han, Y., Cao, J., Lee, S., Ho, K., and An, Z.: Different characteristics of char and soot in the
360 atmosphere and their ratio as an indicator for source identification in Xi’an, China, *Atmos.*
361 *Chem. Phys.*, 10, 595–607, <https://doi.org/10.5194/acp-10-595-2010>, 2010.
- 362 He, H., Wang, Y., Ma, Q., Ma, J., Chu, B., Ji, D., Tang, G., Liu, C., Zhang, H., and Hao, J.:
363 Mineral dust and NO_x promote the conversion of SO₂ to sulfate in heavy pollution days, *Sci.*
364 *Rep.*, 4, 4172, <https://doi.org/10.1038/srep04172>, 2014.
- 365 Huang, J., Wang, T., Wang, W., Li, Z., and Yan, H.: Climate effects of dust aerosols over
366 East Asian arid and semiarid regions, *J. Geophys. Res. Atmos.*, 119, 19,
367 <https://doi.org/10.1002/2014JD021796>, 2014.



- 368 IPCC, 2007: Climate Change 2007: The Physical Science Basis. Contribution of Working
369 Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change
370 [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K. B. Averyt, M. Tignor, and H. L.
371 Miller (eds.)]. Cambridge University Press, Cambridge, United Kingdom and New York, NY,
372 USA, 996 pp.
- 373 IPCC, 2013: Climate Change 2013: The Physical Science Basis. Contribution of Working
374 Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change
375 [Stocker, T. F., D. Qin, G.-K. Plattner, M. Tignor, S. K. Allen, J. Boschung, A. Nauels, Y.
376 Xia, V. Bex, and P. M. Midgley (eds.)]. Cambridge University Press, Cambridge, United
377 Kingdom and New York, NY, USA, 1535 pp.
- 378 Jin, Q. J., Yang, Z. L., and Wei, J. F.: High sensitivity of Indian summer monsoon to Middle
379 East dust absorptive properties, *Sci. Rep.*, 6, 30690, <https://doi.org/10.1038/srep30690>, 2016.
- 380 Kang, Y., Liu, M., Song, Y., Huang, X., Yao, H., Cai, X., Zhang, H., Kang, L., Liu, X., Yan,
381 X., He, H., Zhang, Q., Shao, M., and Zhu, T.: High-resolution ammonia emissions inventories
382 in China from 1980 to 2012, *Atmos. Chem. Phys.*, 16, 2043–2058,
383 <https://doi.org/10.5194/acp-16-2043-2016>, 2016.
- 384 Kim, J., Lee, J., Lee, H. C., Higurashi, A., Take-mura, T., and Song, C. H.: Consistency of the
385 aerosol type classification from satellite remote sensing during the Atmospheric Brown
386 Cloud-East Asia Regional Experiment campaign, *J. Geophys. Res. Atmos.*, 112, D22,
387 <https://doi.org/10.1029/2006JD008201>, 2007.
- 388 Kim, S. B., Yumimoto, K., Uno, I., and Chun, Y.: Dust model intercomparison between
389 ADAM and CFORS/Dust for Asian dust case in 2007 (March 28–April 3), *SOLA*, 7A, 025–
390 028, <https://doi.org/10.2151/sola.7A-007>, 2011.
- 391 Kunwar, B. and Kawamura, K.: One-year observations of carbonaceous and nitrogenous
392 components and major ions in the aerosols from subtropical Okinawa Island, an outflow
393 region of Asian dusts, *Atmos. Chem. Phys.*, 14, 1819–1836, [https://doi.org/10.5194/acp-14-](https://doi.org/10.5194/acp-14-1819-2014)
394 1819-2014, 2014.
- 395 Li, J., Wang, Z., Zhuang, G., Luo, G., Sun, Y., and Wang, Q.: Mixing of Asian mineral dust
396 with anthropogenic pollutants over East Asia: a model case study of a super-dust storm in



- 397 March 2010, Atmos. Chem. Phys., 12, 7591–7607, <https://doi.org/10.5194/acp-12-7591-2012>,
398 2012.
- 399 Lim, S., Lee, M., Lee, G., Kim, S., Yoon, S., and Kang, K.: Ionic and carbonaceous
400 compositions of PM₁₀, PM_{2.5} and PM_{1.0} at Gosan ABC Superstation and their ratios as source
401 signature, Atmos. Chem. Phys., 12, 2007–2024, <https://doi.org/10.5194/acp-12-2007-2012>,
402 2012.
- 403 Lim, S., Lee, M., Kim, S. W., Yoon, S. C., Lee, G., and Lee, Y. J.: Absorption and scattering
404 properties of organic carbon versus sulfate dominant aerosols at Gosan climate observatory in
405 Northeast Asia, Atmos. Chem. Phys., 14, 7781–7793, [https://doi.org/10.5194/acp-14-7781-](https://doi.org/10.5194/acp-14-7781-2014)
406 2014, 2014.
- 407 Liu, D., Abuduwaili, J., Lei, J., Wu, G., and Gui, D.: Wind erosion of saline playa sediments
408 and its ecological effects in Ebinur Lake, Xinjiang, China, Environ. Earth Sci., 63, 241–250,
409 <https://doi.org/10.1007/s12665-010-0690-4>, 2011.
- 410 McFarlane, N. A., Boer, G., Blanchet, J., and Lazare, M.: The Canadian Climate Centre
411 second-generation general circulation model and its equilibrium climate, J. Clim., 5, 1013–
412 1044, [https://doi.org/10.1175/1520-0442\(1992\)005<1013:TCCCSG>2.0.CO;2](https://doi.org/10.1175/1520-0442(1992)005<1013:TCCCSG>2.0.CO;2), 1992.
- 413 Mishra, S. K., Tripathi, S. N., Aggarwal, S. G., and Arola, A.: Effects of particle shape,
414 hematite content and semi-external mixing with carbonaceous components on the optical
415 properties of accumulation mode mineral dust, Atmos. Chem. Phys. Discuss.,
416 <https://doi.org/10.5194/acpd-10-31253-2010>, 2010.
- 417 Nazari, S., Kermani, M., Fazlzadeh, M., Alizadeh-Matboo, S., and Yari, A. R.: The origins
418 and sources of dust particles, their effects on environment and health and control strategies: A
419 review, J. Air Pollut. Health, 1, 137–152, 2016.
- 420 Nie, W., Ding, A., Wang, T., Kerminen, V. M., George, C., Xue, L., Wang, W., Zhang, Q.,
421 Petäjä, T., Qi, X., Gao, X., Wang, X., Yang, X., Fu, C., and Kulmala, M.: Polluted dust
422 promotes new particle formation and growth, Sci. Rep., 4, 6634,
423 <https://doi.org/10.1038/srep06634>, 2014.
- 424 O'Dowd, C. D., Facchini, M. C., Cavalli, F., Ceburnis, D., Mircea, M., Decesari, S., and
425 Putaud, J. P.: Biogenically driven organic contribution to marine aerosol, Nature, 431, 676–
426 680, <https://doi.org/10.1038/nature02959>, 2004.



- 427 Pabortsava, K., Lampitt, R. S., Benson, J., Crowe, C., McLachlan, R., Le Moigne, F. A.,
428 Mark Moore, C., Pebody, C., Provost, P., Rees, A. P., Tilstone, G. H., and Woodward, E.
429 Malcolm S.: Carbon sequestration in the deep Atlantic enhanced by Saharan dust, *Nat.*
430 *Geosci.*, 10, 189–194, <https://doi.org/10.1038/ngeo2899>, 2017.
- 431 Saleh, R., Robinson, E. S., Tkacik, D. S., Ahern, A. T., Liu, S., Aiken, A. C., Sullivan, R. C.,
432 Presto, A. A., Dubey, M. K., Yokelson, R. J., Donahue, N. M., and Robinson, A. L.:
433 Brownness of organics in aerosols from biomass burning linked to their black carbon content,
434 *Nature Geosci.*, advance online publication, 10.1038/ngeo2220,
435 <https://doi.org/10.1038/ngeo2220>, 2014.
- 436 Shang, X., Lee, M., Han, J., Kang, E., Kim, S. W., Gustafsson, Ö., Chang, L.: Identification
437 and Chemical Characteristics of Distinctive Chinese Outflow Plumes Associated with
438 Enhanced Submicron Aerosols at the Gosan Climate Observatory, *Aerosol Air Qual. Res.*, 18:
439 330–342, <https://doi.org/10.4209/aaqr.2017.03.0115>, 2018.
- 440 Shin, S.-K., Müller, D., Lee, C., Lee, K. H., Shin, D., Kim, Y. J., and Noh, Y. M.: Vertical
441 variation of optical properties of mixed Asian dust/pollution plumes according to pathway of
442 air mass transport over East Asia, *Atmos. Chem. Phys.*, 15, 6707–6720,
443 <https://doi.org/10.5194/acp-15-6707-2015>, 2015.
- 444 Su, L. and Toon, O. B.: Saharan and Asian dust: similarities and differences determined by
445 CALIPSO, AERONET, and a coupled climate-aerosol microphysical model, *Atmos. Chem.*
446 *Phys.*, 11, 3263–3280, <https://doi.org/10.5194/acp-11-3263-2011>, 2011.
- 447 Tao, J., Gao, J., Zhang, L., Zhang, R., Che, H., Zhang, Z., Lin, Z., Jing, J., Cao, J., and Hsu,
448 S.-C.: PM_{2.5} pollution in a megacity of southwest China: source apportionment and
449 implication, *Atmos. Chem. Phys. Discuss.*, 14, 5147–5196, [https://doi.org/10.5194/acpd-14-](https://doi.org/10.5194/acpd-14-5147-2014)
450 5147-2014, 2014.
- 451 Textor, C., Schulz, M., Guibert, S., Kinne, S., Balkanski, Y., Bauer, S., Berntsen, T., Berglen,
452 T., Boucher, O., Chin, M., Dentener, F., Diehl, T., Easter, R., Feichter, H., Fillmore, D., Ghan,
453 S., Ginoux, P., Gong, S., Grini, A., Hendricks, J., Horowitz, L., Huang, P., Isaksen, I., Iversen,
454 I., Kloster, S., Koch, D., Kirkevåg, A., Kristjansson, J. E., Krol, M., Lauer, A., Lamarque, J.
455 F., Liu, X., Montanaro, V., Myhre, G., Penner, J., Pitari, G., Reddy, S., Seland, Ø., Stier, P.,
456 Takemura, T., and Tie, X.: Analysis and quantification of the diversities of aerosol life cycles



- 457 within AeroCom, Atmos. Chem. Phys., 6, 1777–1813, <https://doi.org/10.5194/acp-6-1777->
458 2006, 2006.
- 459 Von Salzen, K., McFarlane, N. A., and Lazare, M.: The role of shallow convection in the
460 water and energy cycles of the atmosphere, Clim. Dynam., 25, 671–688,
461 <https://doi.org/10.1007/s00382-005-0051-2>, 2005.
- 462 Warner, J. X., Dickerson, R. R., Wei, Z., Strow, L. L., Wang, Y., and Liang, Q.: Increased
463 atmospheric ammonia over the world's major agricultural areas detected from space, Geophys.
464 Res. Lett., 44, 2875–2884, <https://doi.org/10.1002/2016GL072305>, 2017.
- 465 World Health Organization, and UNAIDS: Air quality guidelines: global update 2005, World
466 Health Organization, 37–41, 2006.
- 467 Yang, M., Howell, S., Zhuang, J., and Huebert, B.: Attribution of aerosol light absorption to
468 black carbon, brown carbon, and dust in China—interpretations of atmospheric measurements
469 during EAST-AIRE, Atmos. Chem. Phys., 9, 2035–2050, <https://doi.org/10.5194/acp-9-2035->
470 2009, 2009.
- 471 Zender, C. S., Miller, R. L., and Tegen, I.: Quantifying mineral dust mass budgets:
472 Terminology, constraints, and current estimates, EOS Trans., 85, 509–512, [https://doi.org/](https://doi.org/10.1029/2004EO480002)
473 10.1029/2004EO480002, 2004.
- 474 Zhang, J.: Coastal saline soil rehabilitation and utilization based on forestry approaches in
475 China, Springer, 11–13, <https://doi.org/10.1007/978-3-642-39915-2>, 2014.
- 476 Zhang, L., Li, Q. B., Gu, Y., Liou, K. N., and Meland, B.: Dust vertical profile impact on
477 global radiative forcing estimation using a coupled chemical-transport–radiative-transfer
478 model, Atmos. Chem. Phys., 13, 7097–7114, <https://doi.org/10.5194/acp-13-7097-2013>,
479 2013a.
- 480 Zhang, R., Jing, J., Tao, J., Hsu, S.-C., Wang, G., Cao, J., Lee, C. S. L., Zhu, L., Chen, Z.,
481 Zhao, Y., and Shen, Z.: Chemical characterization and source apportionment of PM_{2.5} in
482 Beijing: seasonal perspective, Atmos. Chem. Phys., 13, 7053–7074,
483 <https://doi.org/10.5194/acp-13-7053-2013>, 2013b.



- 484 Zhang, R. J., Wang, M. X., Zhang, X. Y., and Zhu, G. H.: Analysis on the Chemical and
485 Physical Properties of Particles in a Dust Storm in Spring in Beijing, *Powder Technol.*, 137,
486 77–82, <https://doi.org/10.1016/j.powtec.2003.08.056>, 2003.
- 487 Zhang, X. L., Wu, G. J., Zhang, C. L., Xu, T. L., and Zhou, Q. Q.: What's the real role of iron-
488 oxides in the optical properties of dust aerosols?, *Atmos. Chem. Phys. Discuss.*, 15, 5619–
489 5662, <https://doi.org/10.5194/acpd-15-5619-2015>, 2015.
- 490 Zhang, X. Y., Zhuang, G. S., Yuan, H., Rahn, K. A., Wang, Z. F., and An, Z. S.: Aerosol
491 particles from dried salt-lakes and saline soils carried on dust storms over Beijing, *Terr.*
492 *Atmos. Ocean. Sci.*, 20, 619–628, [https://doi.org/10.3319/TAO.2008.07.11.03\(A\)](https://doi.org/10.3319/TAO.2008.07.11.03(A)), 2009.
- 493 Zhuang, G., Yi, Z., Duce, R. A., and Brown, P. R.: Link between Iron and Sulfur Suggested
494 by the Detection of Fe (II) in Remote Marine Aerosols, *Nature*, 355, 537–539,
495 <https://doi.org/10.1038/355537a0>, 1992.
- 496



497 Table 1. Mean and mean + σ (standard deviation) concentrations [$\mu\text{g}/\text{m}^3$] of mass and
 498 chemical constituents for PM_{10} , $\text{PM}_{2.5}$, and PM_1 at GCO during 2007 ~ 2012.

499

	PM_{10}		$\text{PM}_{2.5}$		PM_1	
	Mean	Mean+ σ	Mean	Mean+ σ	Mean	Mean+ σ
Mass	30	52	19	32	14	25
Cl^-	0.8	1.7	0.1	0.3	0.1	0.2
NO_3^-	2.1	4.1	1.0	2.2	0.7	1.9
SO_4^{2-}	5.5	9.5	4.4	7.5	4.3	7.5
Na^+	1.2	1.9	0.4	0.7	0.2	0.4
NH_4^+	2.8	4.6	2.4	4.0	2.1	3.6
K^+	0.3	0.5	0.2	0.4	0.2	0.4
Mg^{2+}	0.2	0.3	0.1	0.1	0.02	0.04
Ca^{2+}	0.3	0.7	0.1	0.3	0.1	0.2
OC	4.0	6.6	3.4	5.7	2.6	4.3
OC1	0.1	0.2	0.1	0.2	0.1	0.2
OC2	0.8	1.3	0.8	1.3	0.7	1.1
OC3	1.2	2.0	0.9	1.5	0.7	1.1
OC4	0.9	1.7	0.7	1.3	0.4	0.8
OP	0.9	1.7	0.9	1.6	0.7	1.3
EC	1.5	2.9	1.5	2.7	1.2	2.0
EC1	1.2	2.5	1.1	2.3	0.8	1.5
EC2+EC3	0.3	0.5	0.4	0.6	0.4	0.6



500 Table 2. Intercepts and coefficients for multi-linear regression of PM_{10} , $PM_{2.5}$, and PM_1 mass
501 concentrations using the three principle components (PC1, PC2, and PC3).

502

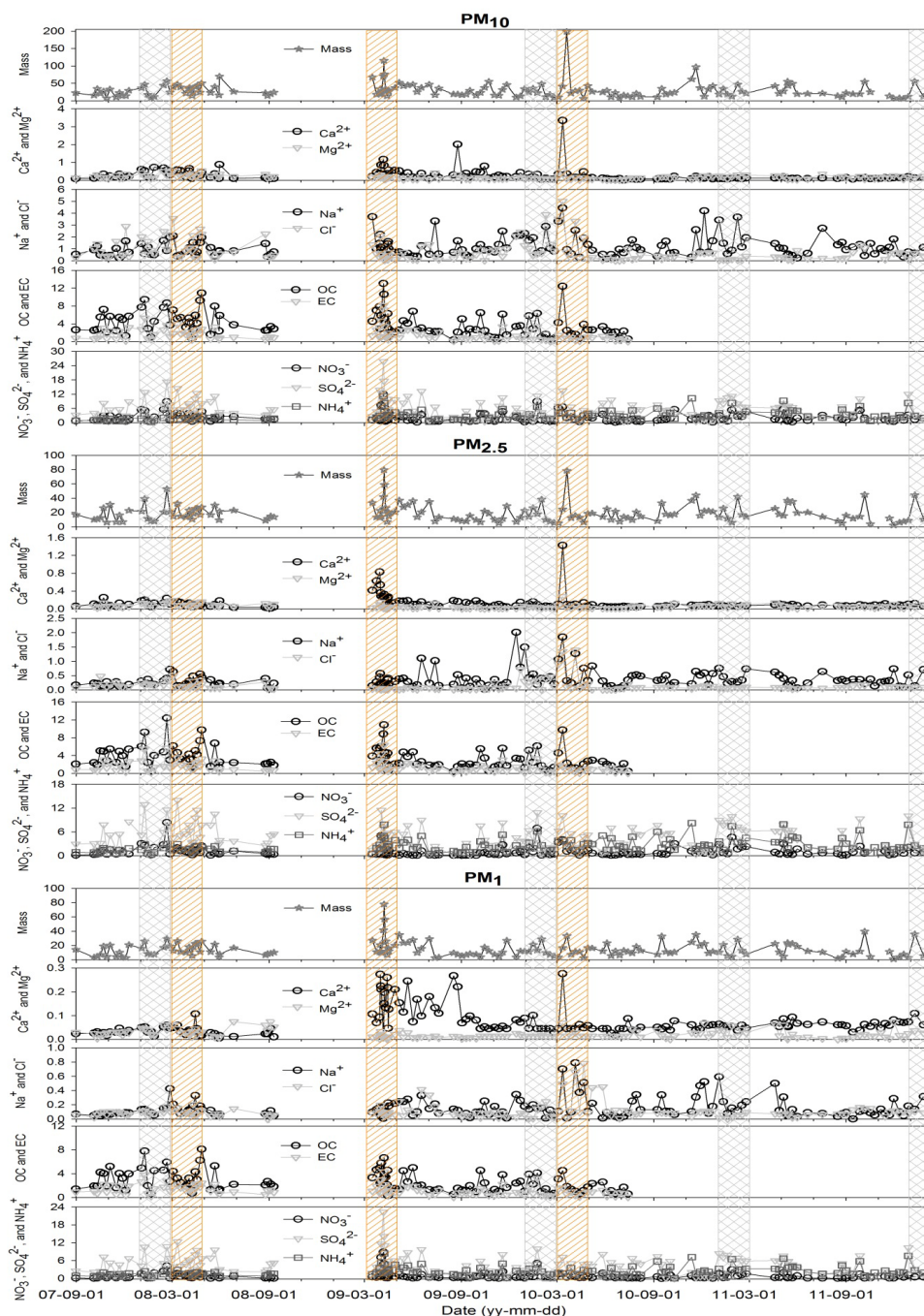
	PM_{10}	$PM_{2.5}$	PM_1
PC1	4.7	3.1	2.5
PC2	3.7	0.4	-0.7
PC3	4.1	1.9	1.6
Intercept	31.1	19.2	14.8



503 Table 3. The statistical summary for mass concentrations of PM₁₀, PM_{2.5}, and PM₁ over entire
504 experiment period [$\mu\text{g}/\text{m}^3$].

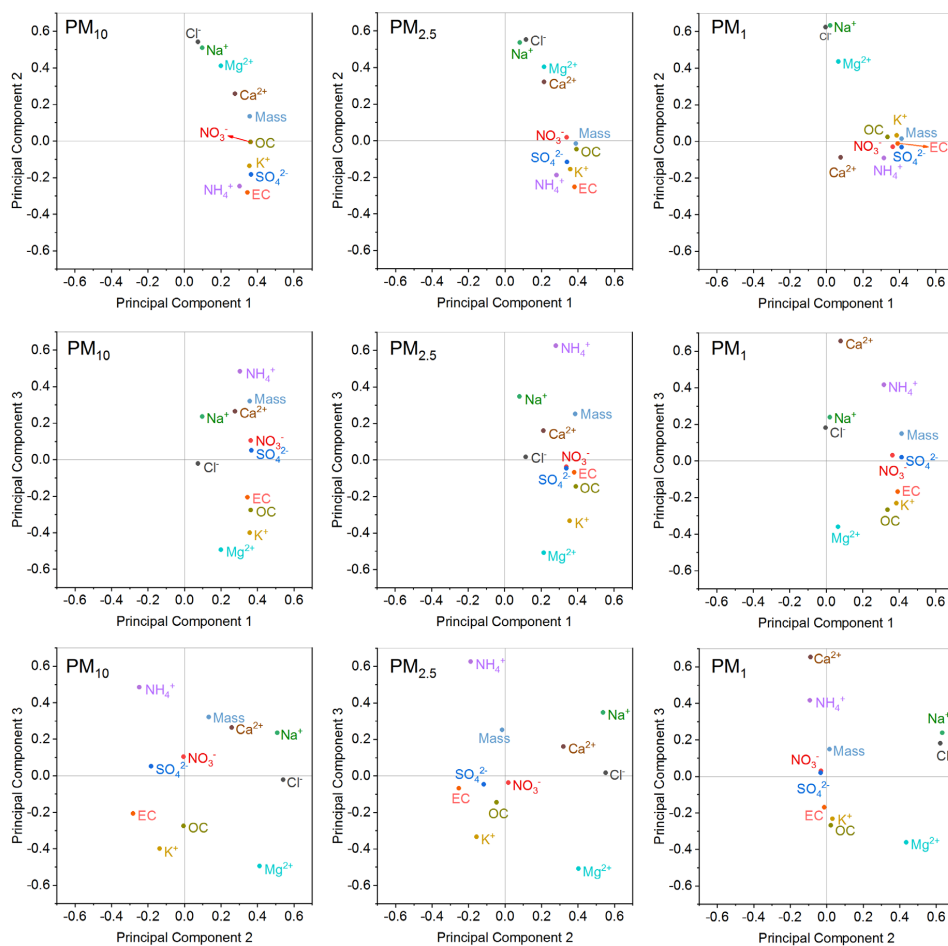
	Median	Mean	S. D. (σ)	Main mode	Mean+ σ
PM ₁₀	24	30	22	25	52
PM _{2.5}	15	19	13	16	32
PM ₁	11	14	11	11	25

505



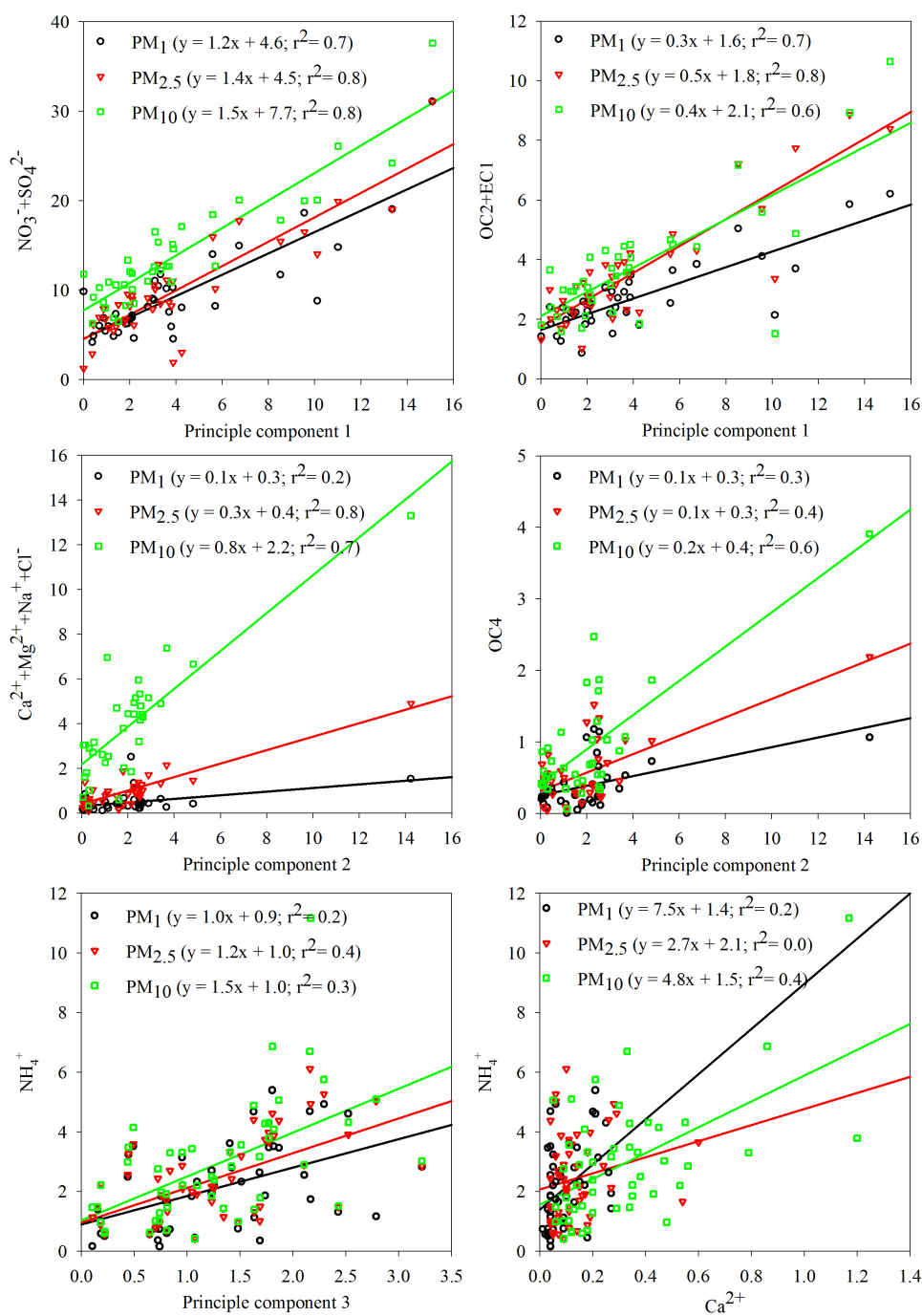
506

507 Figure 1. Time-series variations of major constituents of PM₁₀, PM_{2.5}, and PM₁ for the entire
508 experiment [$\mu\text{g}/\text{m}^3$]. Spring and winter periods are shaded in orange and gray.



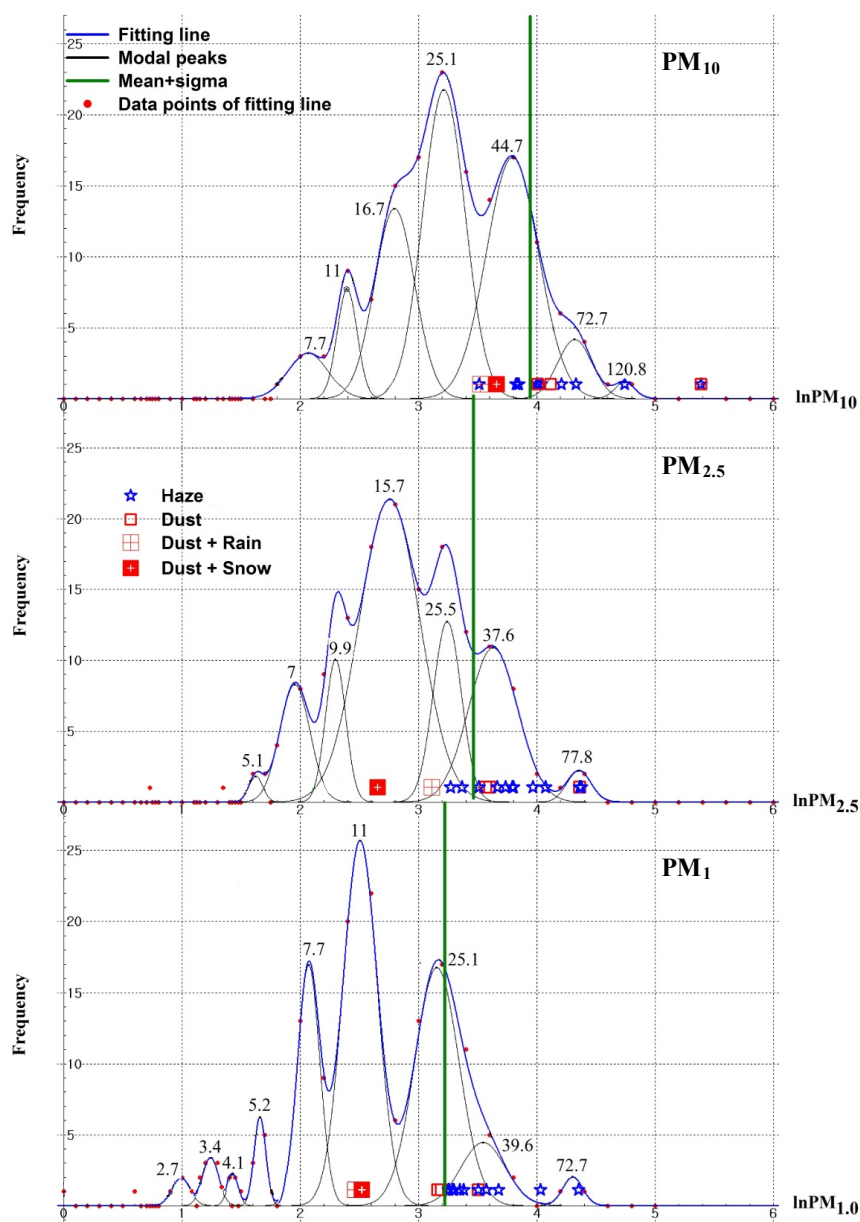
509

510 Figure 2. The results of Principal Component Analysis of all measured species including mass,
511 water-soluble ions, OC, and EC for PM₁₀, PM_{2.5}, and PM₁.



512

513 Figure 3. Correlations between the three principle component loadings and major species
 514 concentrations [$\mu\text{g}/\text{m}^3$] for PM₁₀, PM_{2.5}, and PM₁.



515

516 Figure 4. Frequency distributions of PM_{10} , $PM_{2.5}$ and PM_1 mass concentrations for all
 517 measurements. Mass concentrations are given as \ln values in x-axis. The green
 518 lines stand for $\text{mean} + \sigma$. The individual samples collected during dust or haze
 519 events are marked as different symbols along the x-axis.