

1 Supplementary Materials for

2 **Fine dust emissions from active sands coastal Oceano Dunes, California**

3 This file includes:

- 4 • Supplementary Methods
- 5 • Table S1. Heights of the four optical particle counters (OPCs) above the surface during the Oceano
6 Dunes campaign.
- 7 • Table S2. 35 groups of calibration factors including the best-fit intercept a , uncertainty of intercept σ_a ,
8 slope b , uncertainty of slope σ_b and square of covariance σ_{ab}^2 for the measured aerosol number
9 concentration of each bin of the five optical particle counters (OPCs).
- 10 • Fig. S1. Dust emission from Oceano Dunes during the same saltation event (A) at 2:36 pm on June 2nd,
11 2015 when saltation was weak, and (B) at 4:57 pm on June 2nd, 2015 when saltation was strong.
- 12 • Fig. S2. The experimental configuration of the five optical particle counters (OPCs) at the same height
13 (1.95 meters) above the surface and in a line perpendicular to the wind during June 5th, 6th and 7th, 2015,
14 collecting aerosol mass concentrations for OPC intercalibration.
- 15 • Fig. S3. Calibrated aerosol number concentration against uncalibrated concentration for each of the
16 seven size bins (A) during June 5th to 7th, 2015 and (B) during May 29th to June 4th, 2015.
- 17 • Fig. S4. The ratio of scattered light intensity to incident intensity, $C_{sca} \int P_{11}(\Theta)$, as a function of particle
18 diameter using Lorenz-Mie theory.
- 19 • Fig. S5. Aerosol mass concentration profiles during May 29th to June 4th, 2015 in order of increasing
20 shear velocity.
- 21 • Fig. S6. The normalized volume size distribution of aerosol measured at the four OPC heights during
22 May 29th to June 4th, 2015, (A) when saltation is inactive and (B) when saltation is active.
- 23 • Fig. S7. Vertical aerosol mass flux as a function of shear velocity calculated by using two plans of
24 OPCs' assemblies during May 29th to June 4th, 2015: (A) using D1 and D2 only, and (B) using D3 and
25 D4 only.
- 26 • Fig. S8. Vertical aerosol mass flux (using D1 and D2 only) as a function of shear velocity during May
27 29th to June 4th, 2015 using two regression methods: (A) sea-salt deposition flux is invariant with shear
28 velocity, and (B) sea-salt deposition flux increases non-linearly with shear velocity.
- 29 • Fig. S9. The normalized volume size distribution of dust at emission using the mass flux calculated by
30 the two regression methods.

31

32

33

34 **Supplementary Methods**

35 **Overview**

36 In this document, we present detailed descriptions of the intercalibration and error analysis on the five
37 optical particle counters (OPCs) in Section 1, the correction on the OPC bin sizes using Lorenz-Mie theory
38 in Section 2, the regression on the vertical profiles of aerosol mass concentrations in Section 3, and the
39 methodology to remove the effect of sea-salt deposition flux on these profiles in Section 4.

40 **1. Intercalibration and error analysis on the five optical particle counters (OPCs)**

41 **1.1 Two periods of OPC configurations**

42 We used six identical OPCs (with Series ID 9284, 9290, 9281, 9278, 9287 and 9286, respectively) in the
43 Oceano Dunes field campaign from May 15th to June 7th, 2015. Because OPC 9286 broke on May 25th and
44 therefore could not be included in the subsequent calibration activity, we did not use measurements of this
45 sensor for all of our subsequent analysis. During May 26th to June 4th, 2015, we obtained vertical aerosol
46 number concentrations using four OPCs (of the five good ones) mounted at four different heights at any
47 given time (Table S1 and Fig. 1). During June 5th to June 7th, 2015, we mounted the five good OPCs at the
48 same height and in a line perpendicular to the wind (Fig. S2). The first configuration (May 26th – June 4th)
49 was used to measure fluxes of dust emission, whereas the second configuration (June 5th – 7th) was used to
50 intercalibrate the five OPCs.

51 **1.2 Data-quality control criteria**

52 For the first configuration (May 26th – June 4th), we had three data-quality control criteria. After we built
53 30-minute time blocks and assigned each one second-averaged measurement of aerosol concentration into
54 its corresponding time block, (1) we eliminated those blocks containing fewer than 1800 valid data points
55 (i.e., 30 minutes * 60 second-averaged measurements per minute), (2) we eliminated those blocks with their
56 mean wind directions outside of ± 45 degrees relative to the daily mean wind (Martin and Kok, 2017; Martin
57 et al., 2018), and (3) we eliminated those blocks from May 26th to May 28th, during which the heights of
58 the four OPC sensors were not consistent with the other blocks after May 29th (Table S1) such that these
59 heights were not appropriate for subtracting sea-salt deposition flux in a consistent way. For the second
60 (intercalibration) configuration (June 5th – 7th), we only applied the first quality control criterion, because
61 we were not considered with obtaining vertical profiles for dust flux calculations.

62 **1.3 Intercalibration of the five OPCs during June 5th to 7th**

63 During the second configuration (June 5th to 7th, 2015), the five OPCs were set at the same height to measure
64 aerosol concentrations simultaneously (Fig. S2). For each of the seven size bins (we ignored the largest bin
65 #8 due to insufficient particle counts for constraining uncertainty), we applied linear-least squares
66 regression on aerosol concentration of each of the five OPCs against the mean of the five OPCs, and
67 obtained a correction factor with uncertainty for each OPC (Table S2 and Fig. S3).

68 Specifically, for each of the seven size bins, we used the following equation to obtain the best-fit
 69 intercept and slope for each OPC with respect to the mean value of the five OPCs,

$$70 \quad \ln(y_j) = a_{i,j} + b_{i,j} \cdot \ln(x_{i,j}), \quad (\text{S1.1})$$

71 where the subscripts $i = 1, 2, \dots, 5$ refer to individual OPCs, $j = 1, 2, \dots, 7$ refer to individual size bin, y_j
 72 (in the unit of $\#/m^3$) is the mean concentration of the five OPCs for bin j , and $x_{i,j}$ (in the unit of $\#/m^3$) is
 73 the concentration of each of the five OPCs for bin j . To simplify the following regression procedure, we
 74 define

$$75 \quad u_j = \ln(y_j), \quad (\text{S1.2})$$

$$76 \quad v_{i,j} = \ln(x_{i,j}), \quad (\text{S1.3})$$

77 such that Eq. (S1.1) converts into

$$78 \quad u_j = a_{i,j} + b_{i,j} * v_{i,j} \quad (\text{S1.4}).$$

79 We used the linear-least squares regression to find the best-fit intercept a , slope b , their
 80 uncertainties σ_a and σ_b and covariance σ_{ab} by (Bevington and Robinson, 2003),

$$81 \quad a_{i,j} = \frac{1}{\Delta_{i,j}} \left(\sum_{k=1}^{N_{i,j}} v_{i,j,k}^2 \sum_{k=1}^{N_{i,j}} u_{j,k} - \sum_{k=1}^{N_{i,j}} v_{i,j,k} \sum_{k=1}^{N_{i,j}} v_{i,j,k} u_{j,k} \right), \quad (\text{S1.5})$$

$$82 \quad b_{i,j} = \frac{1}{\Delta_{i,j}} \left(N_{i,j} \sum_{k=1}^{N_{i,j}} v_{i,j,k} u_{j,k} - \sum_{k=1}^{N_{i,j}} v_{i,j,k} \sum_{k=1}^{N_{i,j}} u_{j,k} \right), \quad (\text{S1.6})$$

$$83 \quad \Delta_{i,j} = N_{i,j} \sum_{k=1}^{N_{i,j}} v_{i,j,k}^2 - \left(\sum_{k=1}^{N_{i,j}} v_{i,j,k} \right)^2, \quad (\text{S1.7})$$

$$84 \quad \sigma_{a,i,j}^2 = \frac{\sigma_{i,j}^2}{\Delta_{i,j}} \sum_{k=1}^{N_{i,j}} v_{i,j,k}^2, \quad (\text{S1.8})$$

$$85 \quad \sigma_{b,i,j}^2 = \frac{N_{i,j} \sigma_{i,j}^2}{\Delta_{i,j}}, \quad (\text{S1.9})$$

$$86 \quad \sigma_{ab,i,j}^2 = \sum_{k=1}^{N_{i,j}} \left(\sigma_{i,j}^2 \frac{\partial a_{i,j}}{\partial u_{j,k}} \frac{\partial b_{i,j}}{\partial u_{j,k}} \right), \quad (\text{S1.10})$$

$$87 \quad \sigma_{i,j}^2 = \frac{1}{N_{i,j} - 2} \sum_{k=1}^{N_{i,j}} (u_{j,k} - a_{i,j} - b_{i,j} * v_{i,j,k})^2, \quad (\text{S1.11})$$

$$88 \quad \frac{\partial a_{i,j}}{\partial u_{j,k}} = \frac{1}{\Delta_{i,j}} \left(\sum_{k=1}^{N_{i,j}} v_{i,j,k}^2 - v_{i,j,k} \sum_{k=1}^{N_{i,j}} v_{i,j,k} \right), \quad (\text{S1.12})$$

89
$$\frac{\partial b_{i,j}}{\partial u_{j,k}} = \frac{1}{\Delta_{i,j}} \left(N_{i,j} v_{i,j,k} - \sum_{k=1}^{N_{i,j}} v_{i,j,k} \right) \quad (\text{S1.13}).$$

90 where the subscript $k = 1, 2, \dots, N_{i,j}$ is individual measurement, $N_{i,j}$ is the number of individual
 91 measurements of the j^{th} size bin of the i^{th} OPC, $\sigma_{a,i,j}$ is the uncertainty of the intercept $a_{i,j}$, $\sigma_{b,i,j}$ is the
 92 uncertainty of the slope $b_{i,j}$, $\sigma_{ab,i,j}$ is the uncertainty of the covariance between the intercept $a_{i,j}$ and the
 93 slope $b_{i,j}$, $\sigma_{i,j}^2$ is the estimate of the variance in $u_{j,k}$, the partial derivatives $\frac{\partial a_{i,j}}{\partial u_{j,k}}$ and $\frac{\partial b_{i,j}}{\partial u_{j,k}}$ are the
 94 quantitative sensitivity of the parameters $a_{i,j}$ and $b_{i,j}$ to the value of each individual $v_{i,j,k}$.

95 Using the procedure above, we obtained 35 groups (5 OPCs * 7 bins) of calibration factors including
 96 the best-fit intercept a , slope b , their uncertainties σ_a and σ_b and covariance σ_{ab} (Table S2), based on
 97 which we generated regressed lines (solid lines in Fig. S3). We then derived the regressed uncertainty $\sigma_{fit,i,j}$
 98 at a given point of the regressed line (dashed lines in Fig. S3) from error propagation on Eq. (S1.1) (p. 98-
 99 115 in Bevington and Robinson, 2003),

100
$$\sigma_{fit,i,j} = \sqrt{\sigma_{a,i,j}^2 \left(\frac{\partial y_j}{\partial a_{i,j}} \right)^2 + \sigma_{b,i,j}^2 \left(\frac{\partial y_j}{\partial b_{i,j}} \right)^2 + 2\sigma_{ab,i,j}^2 \left(\frac{\partial y_j}{\partial a_{i,j}} \right) \left(\frac{\partial y_j}{\partial b_{i,j}} \right)} =$$

 101
$$\sqrt{\sigma_{a,i,j}^2 e^{2\ln(y_{j,k})} + \sigma_{b,i,j}^2 \left(e^{\ln(y_{j,k})} \cdot \ln(x_{i,j,k}) \right)^2 + 2\sigma_{ab,i,j}^2 \left(e^{2\ln(y_{j,k})} \cdot \ln(x_{i,j,k}) \right)} \quad (\text{S1.14}).$$

102 1.4 Calibration factors onto measurements during May 29th to June 4th

103 We applied the 35 groups of calibration factors (Table S2) on OPC measurements during May 29th to
 104 June 4th, 2015 by

105
$$y'_{i,j} = e^{a_{i,j} + b_{i,j} \cdot \ln(x'_{i,j})} \quad (\text{S1.15})$$

106 where $y'_{i,j}$ (in the unit of $\#/m^3$) is the calibrated aerosol concentration of the j^{th} bin of the i^{th} OPC during
 107 May 29th to June 4th, $x'_{i,j}$ (in the unit of $\#/m^3$) is the uncalibrated measurement, $a_{i,j}$ and $b_{i,j}$ are best-fit
 108 intercept and slope listed in Table S2. Similar to Eq. (S1.14), the regressed uncertainty was obtained by
 109 error propagation,

110
$$\sigma'_{fit,i,j}$$

 111
$$= \sqrt{\sigma_{a,i,j}^2 e^{2\ln(y'_{i,j,k})} + \sigma_{b,i,j}^2 \left(e^{\ln(y'_{i,j,k})} \cdot \ln(x'_{i,j,k}) \right)^2 + 2\sigma_{ab,i,j}^2 \left(e^{2\ln(y'_{i,j,k})} \cdot \ln(x'_{i,j,k}) \right)} \quad (\text{S1.16})$$

112 where $y'_{i,j,k}$ (in the unit of $\#/m^3$) is the k^{th} calibrated aerosol concentration of the j^{th} bin of the i^{th}
 113 OPC, $x'_{i,j,k}$ (in the unit of $\#/m^3$) is the uncalibrated measurement, $\sigma_{a,i,j}$, $\sigma_{b,i,j}$ and $\sigma_{ab,i,j}^2$ are the
 114 calibration factors listed in Table S2.

115 Using the procedure above, we obtained the calibrated aerosol number concentration of each bin of
 116 the five OPCs in 30-minute intervals during May 29th to June 4th, 2015 by applying the calibration factors
 117 (Table S2) and acquired the regressed uncertainty range by error propagation (Fig. S3).

118 2. OPC bin size correction by Lorenz-Mie theory

119 Because the output size values from the OPCs were those of polystyrene latex spheres (PSLs) following
 120 the international standard ISO 21501-1:2009 (ISO, 2009), the OPCs were not internally calibrated to any
 121 particular dust mineralogy. The optical sizing of dust is sensitive to differences in the refractive index
 122 between dust and PSLs. As such, we corrected the manufacturer-provided bin sizes of PSLs to dust using
 123 Lorenz-Mie theory (Bohren and Huffman, 1983), thereby approximating dust as spherical particles.

124 Mätzler (2002) presented MATLAB code for Lorenz-Mie theory. The input variables include the
 125 particle complex refractive index m , wavelength in the ambient medium λ , and the particle geometric
 126 diameter d . The output variables, related to our calculation, include the scattering efficiency Q_{sca} and the
 127 two scattering amplitude functions $S_1(\theta)$ and $S_2(\theta)$. We applied the following equations on the MATLAB
 128 code outputs to obtain the phase function $P_{11}(\theta)$, the scattering cross section C_{sca} (Eqs. (5.2.111a),
 129 (5.2.112a) and (5.2.94) of Liou, 2002)

$$130 \quad P_{11}(\theta) = \frac{4\pi}{2 \left(\frac{2\pi}{\lambda}\right)^2 C_{sca}} (|S_1(\theta)|^2 + |S_2(\theta)|^2) \quad (S2.1)$$

$$131 \quad C_{sca} = Q_{sca} \cdot \pi \left(\frac{d}{2}\right)^2 \quad (S2.2).$$

132 The phase function quantifies the angular distribution of scattered intensity and the scattering cross section
 133 quantifies the amount of energy scattered from the incident beam by a particle (Liou, 2002). Therefore, the
 134 scattered intensity I_s in Lorenz-Mie theory can be quantified by (Bohren and Huffman, 1983; Liou, 2002),

$$135 \quad I_s(\theta) = I_i \frac{C_{sca} P_{11}(\theta)}{4\pi r^2} \quad (S2.3),$$

136 where r is the distance between the center of the particle to the receiver of the OPC and I_i is the incident
 137 intensity. The OPC measures scattered intensity within the scattering angle $\theta = 90^\circ \pm 60^\circ$ (information
 138 provided by Met One Engineering Department). As such, the scattered intensity measured by OPC is

$$139 \quad I_s(90^\circ \pm 60^\circ) = \frac{C_{sca} I_i}{4\pi r^2} \int_{30^\circ}^{150^\circ} P_{11}(\theta) \sin(\theta) d\theta \quad (S2.4).$$

140 Although we do not know the value of r and I_i , they are constants and do not affect the calculations that
 141 follow. Therefore, we quantify the scattered intensity measured by OPC using

$$142 \quad 4\pi r^2 \frac{I_s(90^\circ \pm 60^\circ)}{I_i} = C_{sca} \int_{30^\circ}^{150^\circ} P_{11}(\theta) \sin(\theta) d\theta \quad (S2.5).$$

Eq. (S2.5) thus establishes the link between $C_{sca} \cdot \text{integral}(P_{11}(\theta))$, the input complex refractive index m , the laser diode wavelength λ used by OPC, and the particle geometric diameter d .

We corrected the OPC bin sizes through two steps. We first input the manufacturer-provided eight bin boundary diameters (seven bins) of PSLs (Table 1, column 1) and their refractive index ($m = 1.59 - 0i$) into Eq. (S2.5), and output the corresponding $C_{sca} \cdot \text{integral}(P_{11}(\theta))$ for each of the eight bin boundary diameters. Second, we input a range of dust refractive indexes (real part $n = 1.53 \pm 0.03$ and imaginary part $k = -10^{-2.5 \pm 0.3}$, summarized in Kok et al., 2017) and an array of dust diameters into Eq. (S2.5). We determined the dust geometric diameters of the eight bin boundaries that produce the same $C_{sca} \cdot \text{integral}(P_{11}(\theta))$ as that of the eight manufacturer-provided PSLs bin boundary diameters calculated in step one (Table 1, column 2 and Fig. S4).

3. Concentration unit conversion and profile fitting

We generated the concentration profile in preparation for the dust flux calculation using the gradient method (Gillette et al., 1972; Shao, 2008). To obtain the concentration profile, we first converted the calibrated 30-minute aerosol number concentration at four different heights into mass concentrations, after which we fit regressions on the height-resolved mass concentrations.

We converted between the number concentration C_N ($\#/m^3$) and the mass concentration C_M (kg/m^3) as (Seinfeld and Pandis, 2016)

$$C_{M_{i,j}} = \frac{\rho}{6} \pi C_{N_{i,j}} (D_j \cdot 10^{-6})^3, \quad (S3.1)$$

where the subscripts $i = 1, 2, \dots, 4$ refer to the OPCs, $j = 1, 2, \dots, 7$ refer to the size bins, $\rho = 2.5 \pm 0.2 \times 10^3$ kg/m^3 is the dust density (Kok et al., 2017), and D_j (in the unit of μm) is the dust geometric diameters of the j^{th} bin (Table 1, column 2).

In order to obtain the size-resolved mass fluxes, we applied linear-least squares regression (Eq. (3)) to vertical mass concentration profiles for each of the seven size bins. As such, we obtained the concentration profile fits for each size bin associated with concurrent shear velocity measurements (in units of m/s) (Fig. S5).

4. Removal of sea-salt aerosol deposition signal from dust flux calculations

We found that the concentration profiles (Fig. S5) deviated from the logarithmic profiles expected to occur from an active emission source (Stull, 1988; Kind, 1992; Gillies and Berkofsky, 2004), a result that we inferred as the influence of sea-salt aerosol. Because we measured dust concentrations ~ 650 meters from the shoreline, we expect increasing sea-salt aerosol concentration with height due to the upwind deposition of near-surface sea-salt aerosol (Liang et al., 2016). We indeed generally observed an increasing concentration with height for the lowest two or three OPCs when saltation was inactive (horizontal saltation flux $Q = 0$), consistent with sea-salt deposition, but found a decrease in concentration with height when

176 saltation was active ($Q > 0$), consistent with dust emission (Fig. S5). In this section, we describe (1) how
177 we diagnosed whether sea-salt aerosols played a role and (2) how we removed their effect on our
178 measurements to the extent possible.

179 4.1 Diagnosis of sea-salt aerosols deposition

180 We calculated the normalized volume size distribution of the four OPC heights (Table S1) during May
181 29th to June 4th, 2015 separately for times when saltation is inactive (Fig. S6A) and active (Fig. S6B). The
182 aerosol PSD when saltation is inactive is largely coarser than when it is active for all four heights. Because
183 sea-salt aerosols are significantly coarser than dust (O’Dowd and de Leeuw, 2007), we concluded that sea-
184 salt deposition affected our measurements at all four heights.

185 Because lower sensors more strongly reflect local dust emission whereas the higher sensors reflect the
186 upwind signal from the sea-salt deposition, we used partial of the four OPCs to study the contribution of
187 sea-salt aerosols deposition to the dust emission measurements. Specifically, we studied two types of
188 measured concentration profiles to calculate the vertical mass flux: (1) using only D1 and D2 and (2) only
189 D3 and D4 (detailed heights of D1, D2, D3 and D4 are listed in Table S1).

190 Because deposition of sea-salt aerosols leads to negative vertical mass flux ($F_d < 0$) and emission of
191 dust aerosols leads to positive mass flux ($F_d > 0$), we used the difference in the sign of mass flux to
192 distinguish the contributions of sea-salt aerosols versus dust aerosols. Specifically, for each of the two plans
193 mentioned in the previous paragraph, we categorized the gradient-method calculated mass fluxes (Eq. (3))
194 into four scenarios: (1) active saltation and positive vertical mass flux ($Q > 0, F_d > 0$), (2) active saltation
195 and negative vertical mass flux ($Q > 0, F_d < 0$), (3) inactive saltation and positive vertical mass flux ($Q =$
196 $0, F_d > 0$) and (4) inactive saltation and negative mass flux ($Q = 0, F_d < 0$).

197 Fig. S7 shows the results of the four scenarios for the two plans. Using only the lowest two OPCs, D1
198 and D2, we found the aerosol flux to be small and negative (deposition) when saltation was inactive, and
199 large and positive (emission) when saltation was active. In contrast, use of the higher OPCs, D3 and D4,
200 showed a positive aerosol flux when saltation was inactive, which indicated a large dust signal pollution by
201 sea-salt aerosols. This comparison in the two plans supports that lower sensors more strongly reflect local
202 dust emission whereas the higher sensors reflect the upwind signal from the sea-salt deposition (refer to
203 saltation layer height detailed in Martin and Kok, 2017). As such, we used only D1 and D2 to calculate the
204 dust emission flux.

205 4.2 Two regression methods to remove sea-salt deposition flux

206 Because using only the lowest two OPCs did not eliminate the deposition flux of sea-salt from our
207 results, we subtracted mass fluxes measured by D1 and D2 during inactive saltation events (inferred as the
208 background sea-salt deposition signal) from the mass fluxes calculated by D1 and D2 during active saltation
209 events. Since we were unable to find independent detailed *in situ* measurements relating near-shore sea-salt

210 deposition flux to shear velocity (u_*), we considered two different empirical regression methods to correct
211 for the background sea-salt deposition signal: (1) sea-salt deposition flux is invariant with increasing shear
212 velocity and (2) sea-salt deposition flux increases non-linearly with increasing shear velocity. Fig. S8 shows
213 the results of the two regression methods.

214 After we obtained the regressed absolute value of the sea-salt deposition flux during active saltation
215 (red regression lines in Fig. S8), we added the regressed deposition fluxes (i.e., subtracted the equivalent
216 emission values) from the calculated mass fluxes during active saltation (red closed left-pointing triangles
217 in Fig. S8). The summed values are the net dust emission fluxes during active saltation events.

218 We compared the PSD of emitted dust obtained by the two regression methods (Fig. S9). Because the
219 second regression method leads to a decreasing trend of mass flux with increasing shear velocity (Bin #2,
220 3, and 4 in Fig. S8B) and the PSDs of emitted dust obtained by the two regression methods are highly
221 similar (Fig. S9), we chose to use the simpler correction assuming a constant sea-salt deposition flux for
222 obtaining the net dust deposition fluxes presented in the main text.

223

224 5. References:

225 Bevington, P. R. and Robinson, D. K.: Data reduction and error analysis for the physical sciences, 3rd ed.,
226 McGraw-Hill Education., 2003.

227 Bohren, C. F. and Huffman, D. R.: Absorption and scattering of light by small particles, 1st ed., Wiley-
228 VCH., 1983.

229 Gillette, D. A., Blifford, I. H. and Fenster, C. R.: Measurements of aerosol size distributions and vertical
230 fluxes of aerosols on land subject to wind erosion, *J. Appl. Meteorol. Climatol.*, 11(6), 977–987,
231 doi:10.1175/1520-0450(1972)011<0977:MOASDA>2.0.CO;2, 1972.

232 Gillies, J. A. and Berkofsky, L.: Eolian suspension above the saltation layer, the concentration profile, *J.*
233 *Sediment. Res.*, 74(2), 176–183, doi:10.1306/091303740176, 2004.

234 ISO: Determination of particle size distribution: single particle light interaction methods, part 1: Light
235 scattering aerosol spectrometer, ISO 21501-1, 2009.

236 Kind, R. J.: One-dimensional aeolian suspension above beds of loose particles-A new concentration-profile
237 equation, *Atmos. Environ.*, 26A(5), 927–931, doi:10.1016/0960-1686(92)90250-O, 1992.

238 Kok, J. F., Ridley, D. A., Zhou, Q., Miller, R. L., Zhao, C., Heald, C. L., Ward, D. S., Albani, S. and
239 Haustein, K.: Smaller desert dust cooling effect estimated from analysis of dust size and abundance,
240 *Nat. Geosci.*, 10(March), doi:10.1038/ngeo2912, 2017.

241 Liang, T., Chamecki, M. and Yu, X.: Sea salt aerosol deposition in the coastal zone: A large eddy simulation
242 study, *Atmos. Res.*, 180, 119–127, doi:10.1016/j.atmosres.2016.05.019, 2016.

243 Liou, K.-N.: An introduction to atmospheric radiation, 2nd ed., Academic Press, Inc., 2002.

244 Martin, R. L., Kok, J. F., Hugenholtz, C. H., Barchyn, T. E., Chamecki, M. and Ellis, J. T.: High-frequency
245 measurements of aeolian saltation flux: Field-based methodology and applications, *Aeolian Res.*, 30,
246 97–114, doi:10.1016/j.aeolia.2017.12.003, 2018.

247 Martin, R. L. and Kok, J. F.: Wind-invariant saltation heights imply linear scaling of aeolian saltation flux
248 with shear stress, *Sci. Adv.*, 3(6), 1–30, doi:10.1126/sciadv.1602569 ., 2017.

249 Mätzler, C.: MATLAB Functions for Mie Scattering and Absorption, version 2, IAP Res. Rep., 1–24,
250 doi:10.1039/b811392k, 2002.

251 O’Dowd, C. D. and de Leeuw, G.: Marine aerosol production: a review of the current knowledge, *Philos.*
252 *Trans. R. Soc. A Math. Phys. Eng. Sci.*, 365(1856), 1753–1774, doi:10.1098/rsta.2007.2043, 2007.

253 Seinfeld, J. H. and Pandis, S. N.: *Atmospheric chemistry and physics: from air pollution to climate change*,
254 3rd ed., Wiley., 2016.

255 Shao, Y.: *Physics and modelling of wind erosion*, 2nd ed., edited by Y. Shao, Springer, Netherlands., 2008.

256 Stull, R. B.: *An introduction to boundary layer meteorology*, 1st ed., Springer Netherlands., 1988.

257
258
259

260 **Tables and Figures**

261 Table S1. Heights of the four optical particle counters (OPCs) above the surface during the Oceano Dunes
262 campaign. D1 denotes the OPC set at the lowest height, D2 the second lowest, D3 the third lowest and D4
263 at the highest height (see Fig. 1).

Measurement Time (Pacific Time) (Month-Day Start Time- End Time)		D1 (m)	D2 (m)	D3 (m)	D4 (m)
05-15 13:03- 18:14		1.62	3.54	6.44	None
05-16 10:48- 17:47		1.62	3.54	6.44	None
05-19 13:50- 15:18		1.62	2.74	3.54	None
05-23 12:24- 18:28		1.62	2.74	3.54	6.44
05-24	11:27- 12:37	1.62	2.74	3.54	6.44
	13:08- 16:40	1.62	2.74	6.44	None
05-26 13:07- 18:01		1.11	1.62	2.74	6.44
05-27 12:04- 18:01		1.11	1.62	2.74	6.44
05-28 12:20- 17:32		1.11	1.62	2.74	6.44
05-29 11:22- 17:08		0.74	1.48	2.29	6.44
05-30 10:16- 18:16		0.74	1.48	2.29	6.44
05-31 10:03- 18:01		0.74	1.48	2.29	6.44
06-01 10:26- 15:35		0.74	1.48	2.29	6.44
06-02 10:47- 18:14		0.75	1.49	2.31	6.44
06-03 11:57- 18:36		0.75	1.49	2.30	6.44
06-04 11:24- 15:22		0.75	1.49	2.31	6.44

264

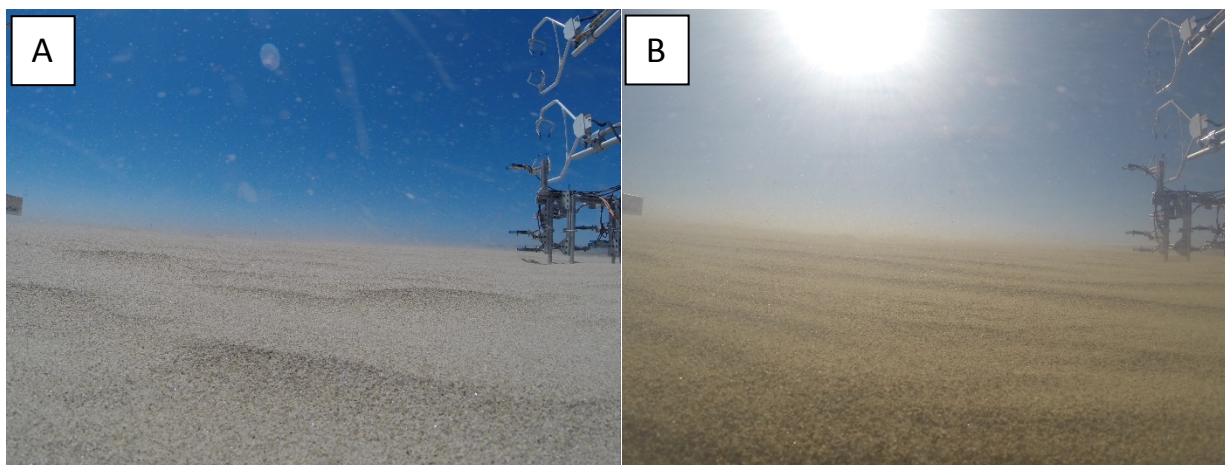
265

266

267 Table S2. 35 groups of calibration factors including the best-fit intercept a , uncertainty of intercept σ_a ,
268 slope b , uncertainty of slope σ_b and square of covariance σ_{ab}^2 for the measured aerosol number
269 concentration of each bin of the five optical particle counters (OPCs). These regression constants are
270 obtained using the linear-least squares regression method.

OPC Series ID		$a \pm \sigma_a$	$b \pm \sigma_b$	σ_{ab}^2
9284	Bin_1	0.1485 \pm 0.3228	0.9850 \pm 0.0511	-0.0165
	Bin_2	0.0463 \pm 0.4775	1.0110 \pm 0.0811	-0.0387
	Bin_3	-0.1671 \pm 0.6240	1.0506 \pm 0.1112	-0.0693
	Bin_4	0.0320 \pm 0.6623	1.0146 \pm 0.1292	-0.0855
	Bin_5	0.4924 \pm 0.5089	0.9251 \pm 0.1089	-0.0554
	Bin_6	1.3569 \pm 0.4838	0.7427 \pm 0.1324	-0.0639
	Bin_7	1.9912 \pm 0.3033	0.5138 \pm 0.1138	-0.0342
9290	Bin_1	0.1391 \pm 0.3233	0.9799 \pm 0.0508	-0.0164

	Bin_2	0.2129 ± 0.4656	0.9709 ± 0.0782	-0.0364
	Bin_3	0.2331 ± 0.5881	0.9693 ± 0.1037	-0.0609
	Bin_4	0.1176 ± 0.6682	0.9853 ± 0.1287	-0.0860
	Bin_5	0.0669 ± 0.5774	0.9881 ± 0.1202	-0.0694
	Bin_6	0.4025 ± 0.7106	0.9079 ± 0.1760	-0.1249
	Bin_7	0.6305 ± 0.5245	0.8126 ± 0.1563	-0.0818
9281	Bin_1	0.1524 ± 0.3216	0.9965 ± 0.0515	-0.0166
	Bin_2	0.0752 ± 0.4729	1.0190 ± 0.0814	-0.0385
	Bin_3	-0.0575 ± 0.6124	1.0569 ± 0.1118	-0.0685
	Bin_4	0.4705 ± 0.6113	0.9506 ± 0.1220	-0.0746
	Bin_5	0.7238 ± 0.4814	0.8861 ± 0.1043	-0.0502
	Bin_6	1.3071 ± 0.4939	0.7385 ± 0.1320	-0.0651
	Bin_7	1.8436 ± 0.2680	0.5393 ± 0.0950	-0.0252
9278	Bin_1	-0.1650 ± 0.3384	1.0092 ± 0.0522	-0.0177
	Bin_2	-0.1400 ± 0.4896	0.9987 ± 0.0797	-0.0390
	Bin_3	-0.0487 ± 0.6072	0.9759 ± 0.1025	-0.0622
	Bin_4	-0.0345 ± 0.6741	0.9683 ± 0.1239	-0.0835
	Bin_5	-0.2932 ± 0.6063	1.0223 ± 0.1214	-0.0736
	Bin_6	-0.4668 ± 0.8268	1.0350 ± 0.1887	-0.1559
	Bin_7	-0.3394 ± 0.6523	0.9947 ± 0.1756	-0.1143
9287	Bin_1	-0.0045 ± 0.3301	0.9928 ± 0.0514	-0.0170
	Bin_2	0.1136 ± 0.4699	0.9656 ± 0.0771	-0.0362
	Bin_3	0.2728 ± 0.5741	0.9369 ± 0.0985	-0.0565
	Bin_4	0.2557 ± 0.6319	0.9441 ± 0.1199	-0.0757
	Bin_5	0.3338 ± 0.5252	0.9201 ± 0.1078	-0.0566
	Bin_6	0.1567 ± 0.7051	0.9496 ± 0.1711	-0.1205
	Bin_7	0.2439 ± 0.5531	0.9123 ± 0.1621	-0.0894



271
272 Figure S1. Dust emission from Oceano Dunes during the same saltation event (A) at 2:36 pm on June 2nd,
273 2015 when saltation was weak, and (B) at 4:57 pm on June 2nd, 2015 when saltation was strong. The two
274 sonic anemometers in the top right corner are located at 0.64 meters and 1.16 meters above the surface.
275 Although sun glare obscures (B), the increase in haziness in the upper half of the photograph suggests active
276 dust emission from the study site.

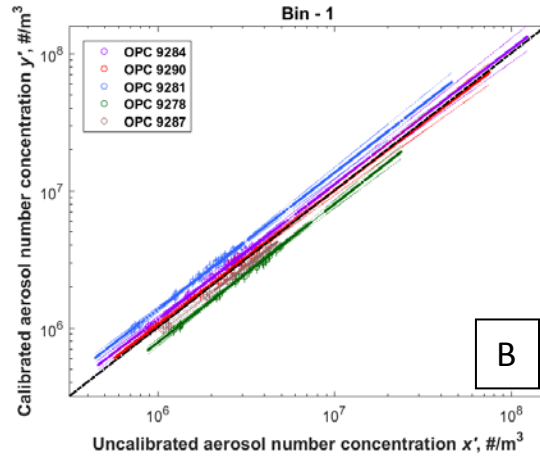
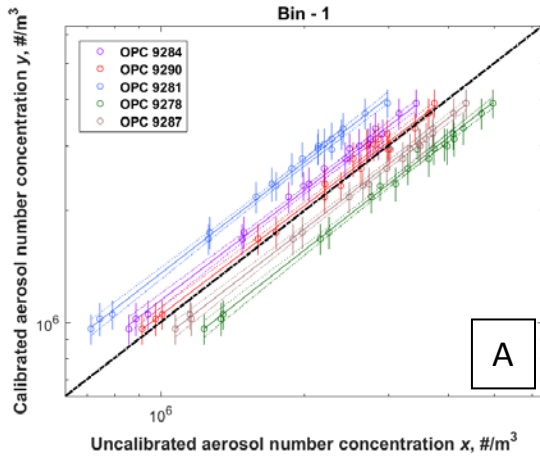
277
278



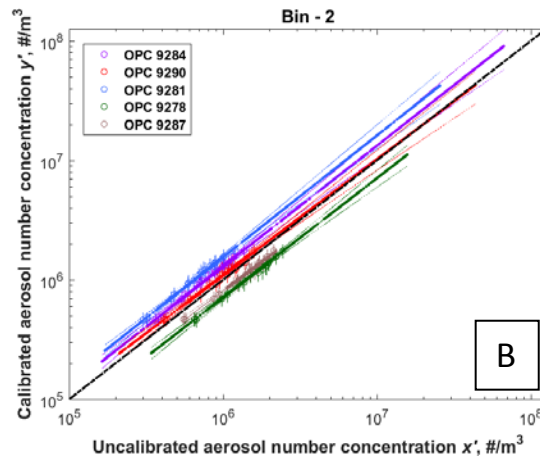
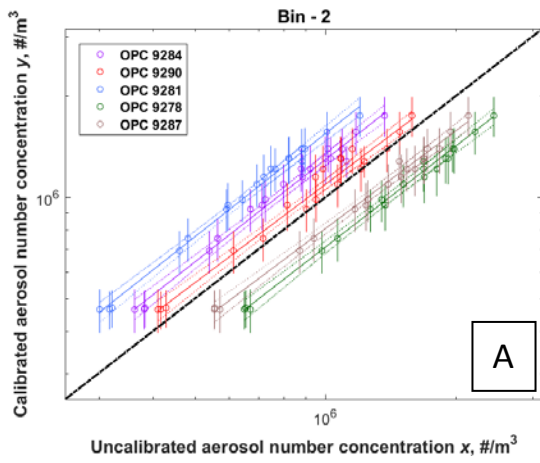
279
280 Figure S2. The experimental configuration of the five optical particle counters (OPCs) (with Series ID 9284,
281 9290, 9281, 9278 and 9287, respectively) at the same height (1.95 meters) above the surface and in a line
282 perpendicular to the wind during June 5th, 6th and 7th, 2015, collecting aerosol number concentrations for
283 OPC intercalibration.

284
285

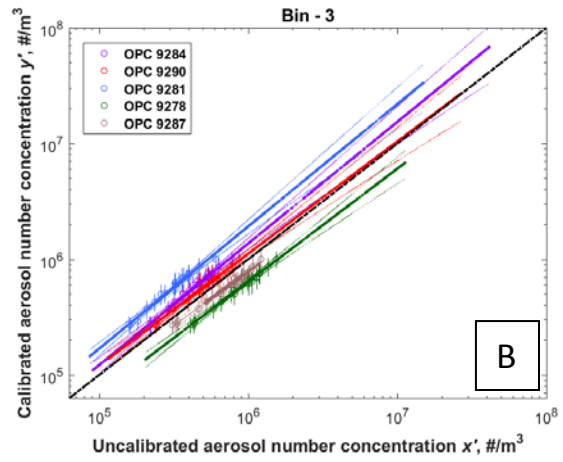
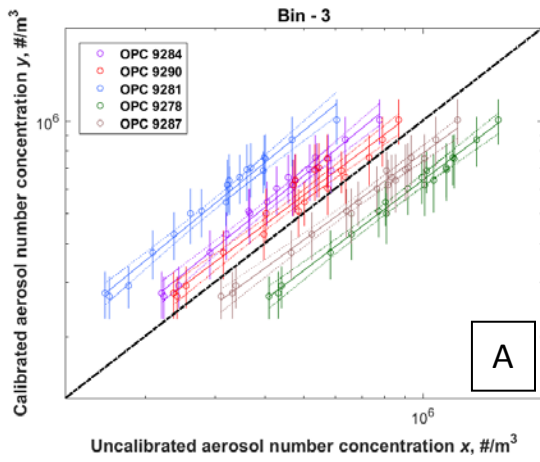
286



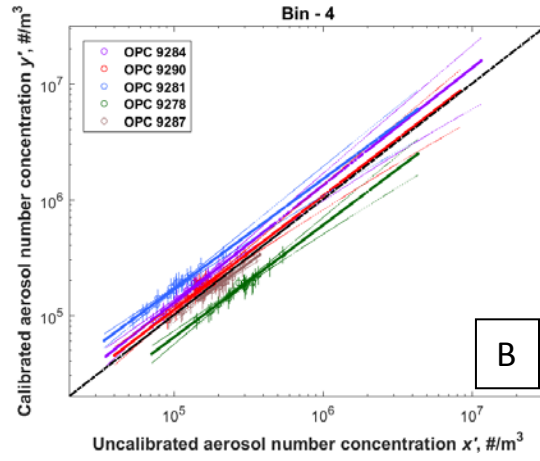
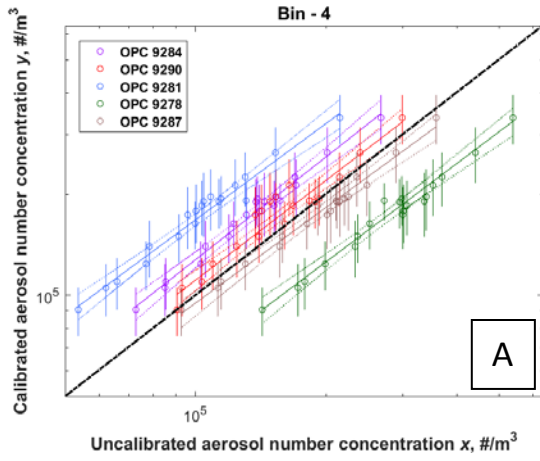
287



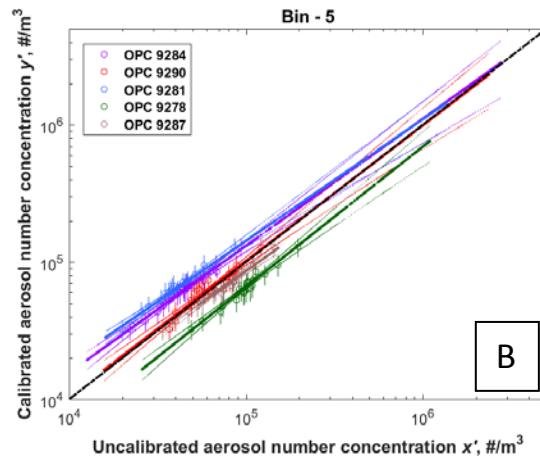
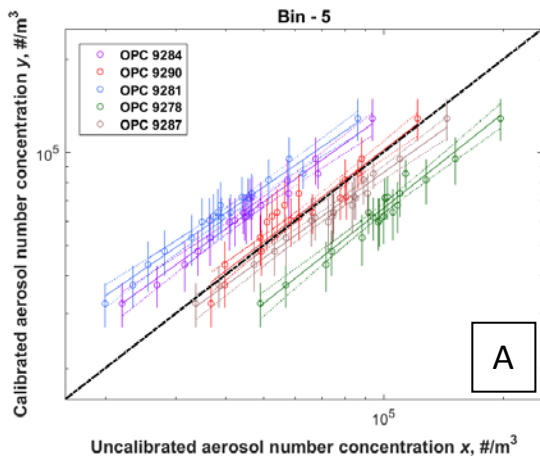
288



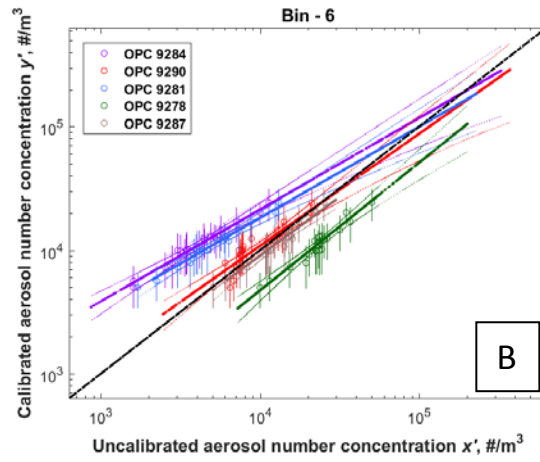
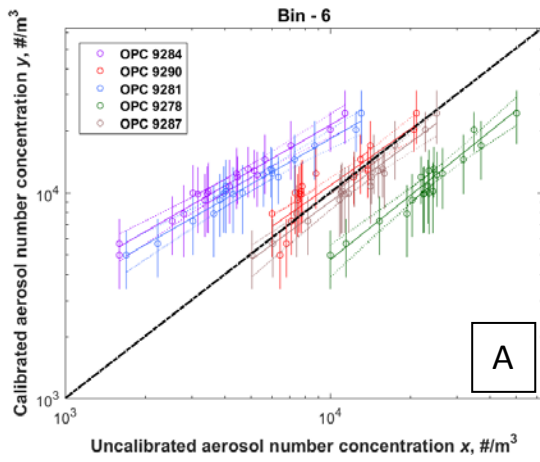
289

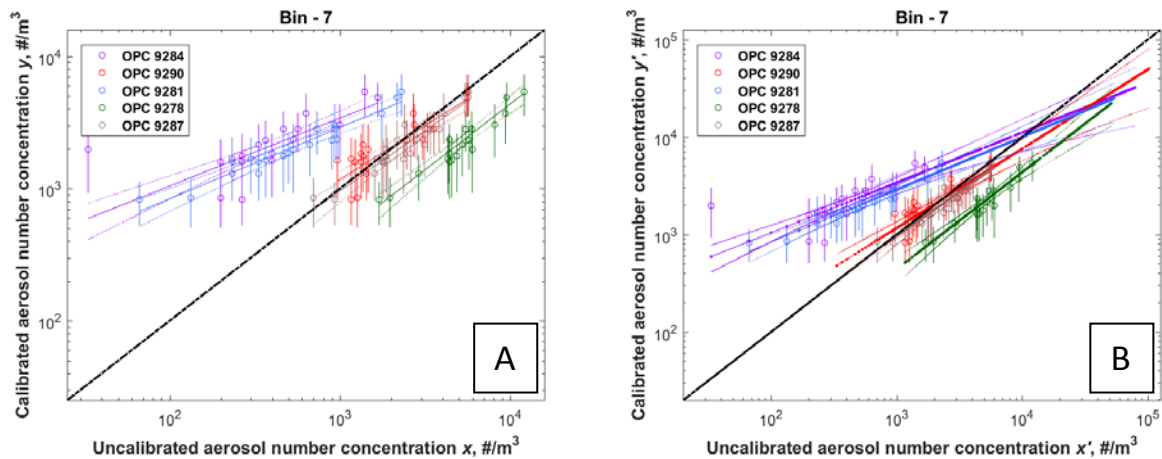


290

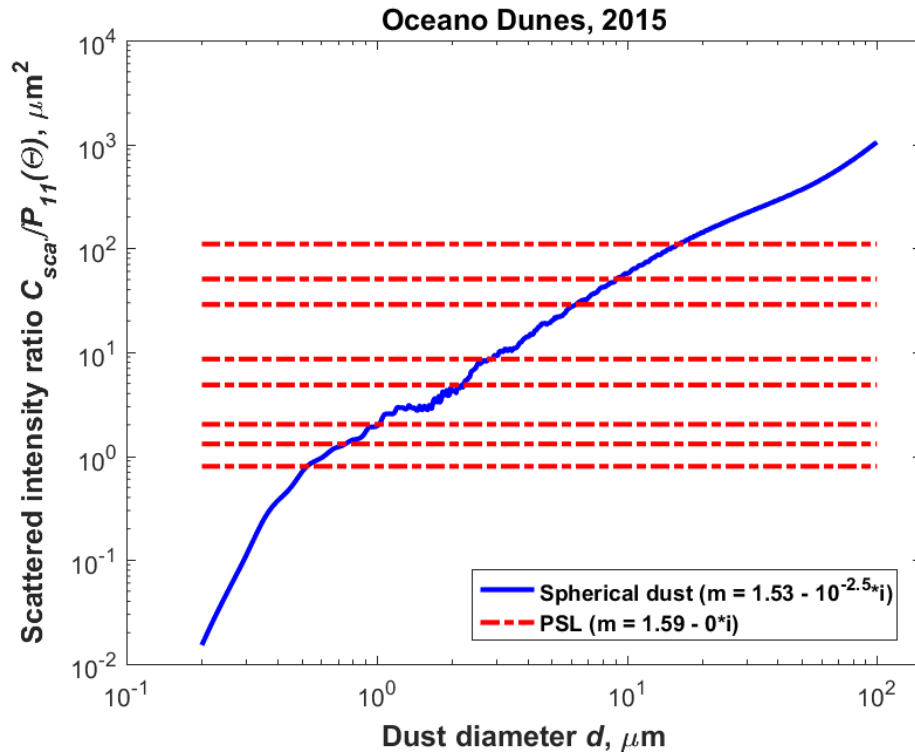


291



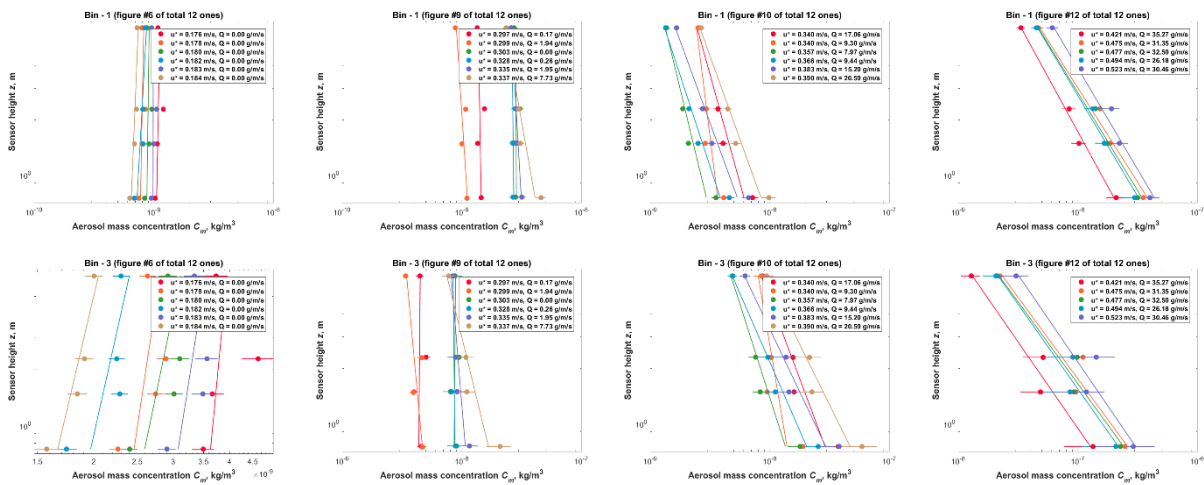


292
 293 Figure S3. Calibrated aerosol number concentration against uncalibrated concentration for each of the seven
 294 size bins (A) during June 5th to 7th, 2015 and (B) during May 29th to June 4th, 2015. Each of the 14 plots
 295 includes the mean number concentration of the five OPCs against concentration measured by each of the
 296 five OPCs during June 5th, 6th and 7th (open circles), standard error of the concentration by the five OPCs
 297 during June 5th, 6th and 7th (error bars), linear-least squares regression lines (solid lines), regressed
 298 uncertainty range within ± 1 standard deviation (dashed lines), and 1:1 reference line (black dot-dashed
 299 line). The seven plots of (B) also include the regressed concentration of each of the five individual OPCs
 300 against uncalibrated concentration measured by each of the five individual OPCs during May 29th to June
 301 4th (small closed circles).
 302



303
 304 Figure S4. The ratio of scattered light intensity to incident intensity, $C_{sca} \int P_{11}(\Theta)$, as a function of particle
 305 diameter using Lorenz-Mie theory. The figure includes the ratios for each of the eight polystyrene latex
 306 spheres (PSLs) boundary diameter sizes (red dash-dot lines) (Table 1, column 1) and the ratios for dust
 307 particles with various geometric diameters (blue line).

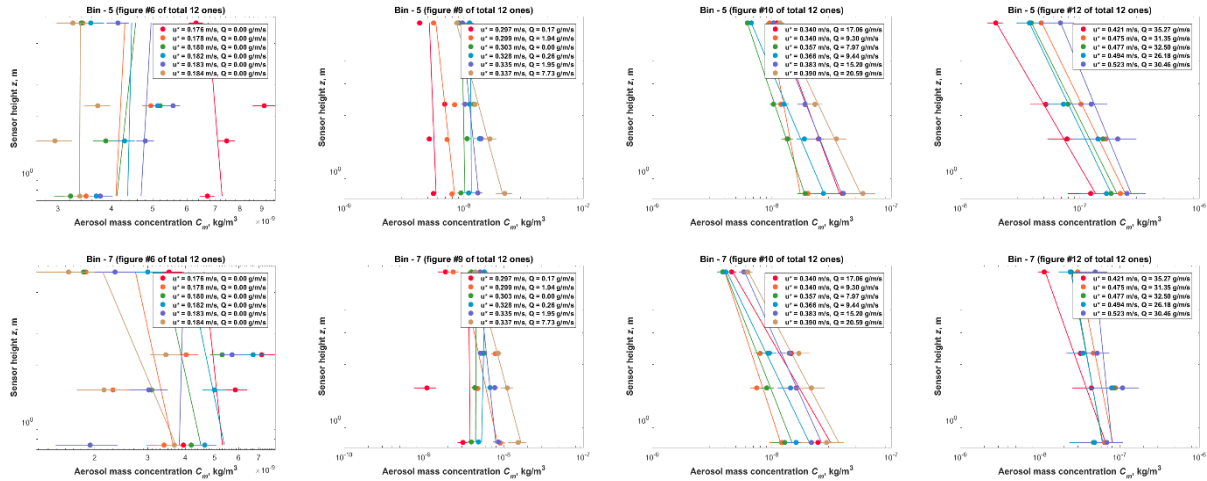
308
 309



310

311

312

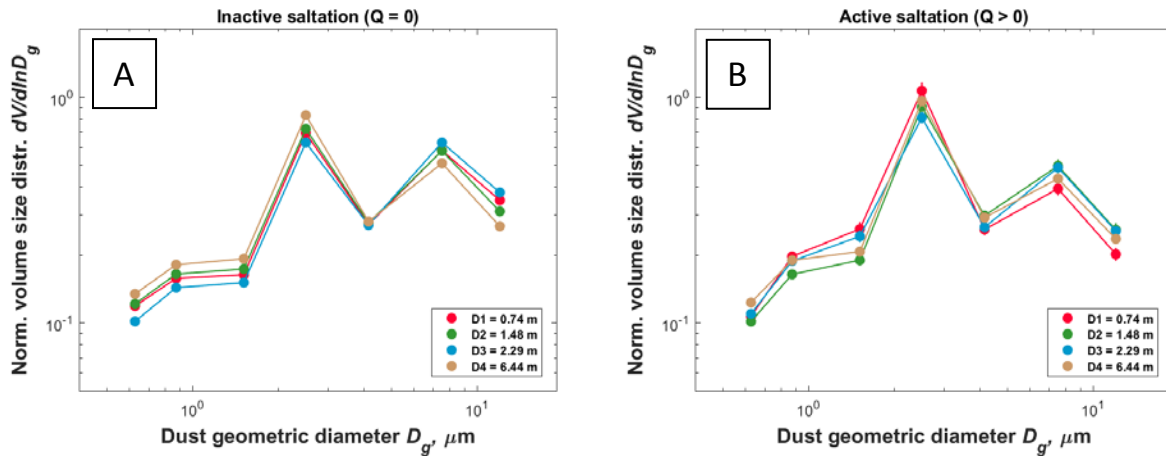


313

314 Figure S5. Aerosol mass concentration profiles during May 29th to June 4th, 2015 in order of increasing
 315 shear velocity. Each of the 84 subplots (12 subplots * 7 bins) includes the aerosol mass concentrations at
 316 four heights (close circles), uncertainties of the mass concentrations (error bars), and linear-least squares
 317 regression lines (lines). The value of shear velocity u_* (in the unit of m/s) and the value of horizontal
 318 saltation flux Q (in the unit of g/m/s) are provided in the legend. Note that we only present 16 of the 84
 319 subplots due to the limitation in space.

320

321

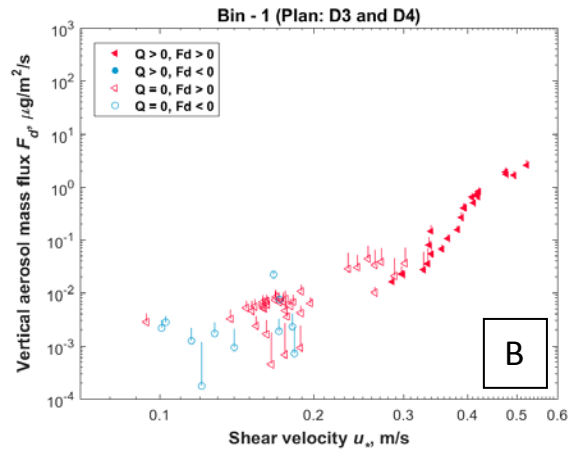
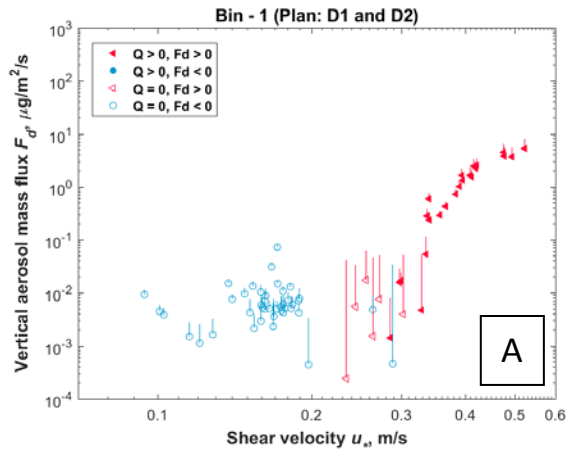


322

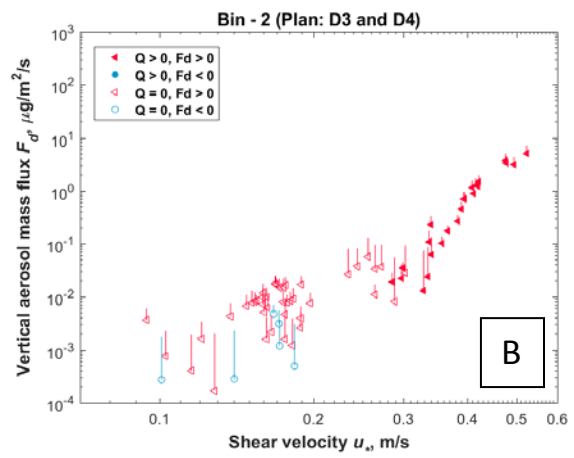
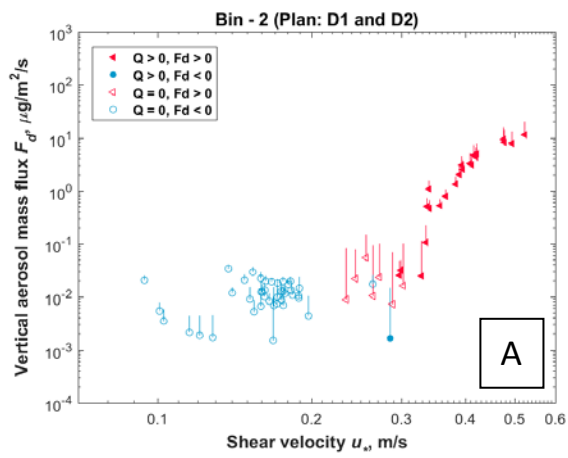
323 Figure S6. The normalized volume size distribution of aerosol measured at the four OPC heights during
 324 May 29th to June 4th, 2015, (A) when saltation is inactive (horizontal saltation flux $Q = 0$) and (B) when
 325 saltation is active ($Q > 0$).

326

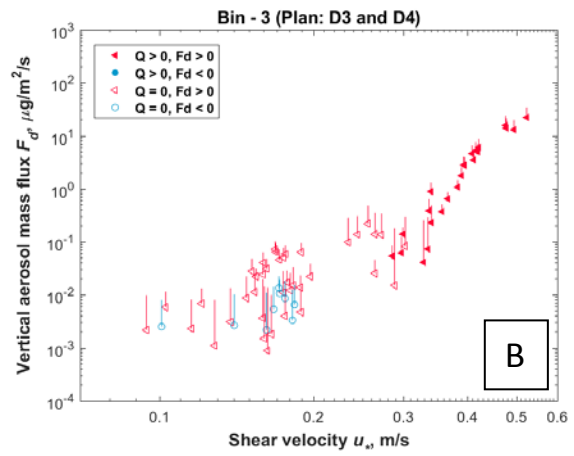
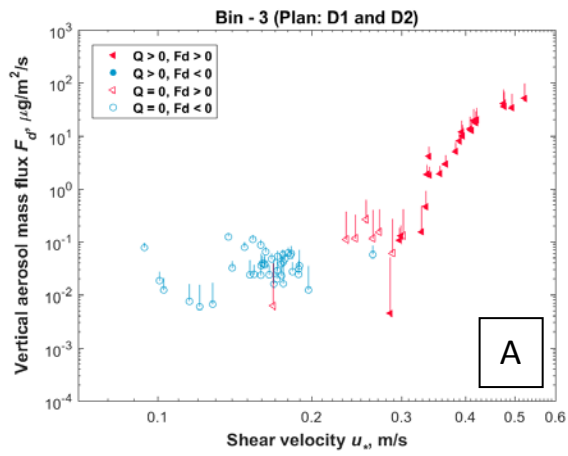
327



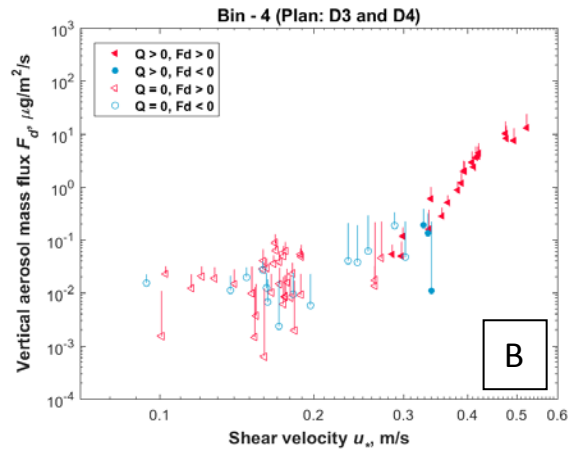
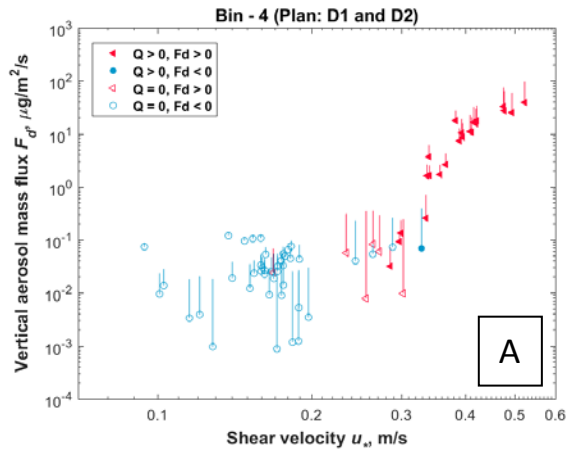
328



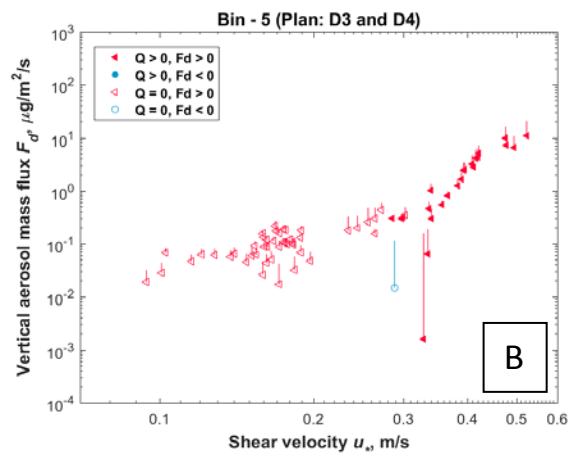
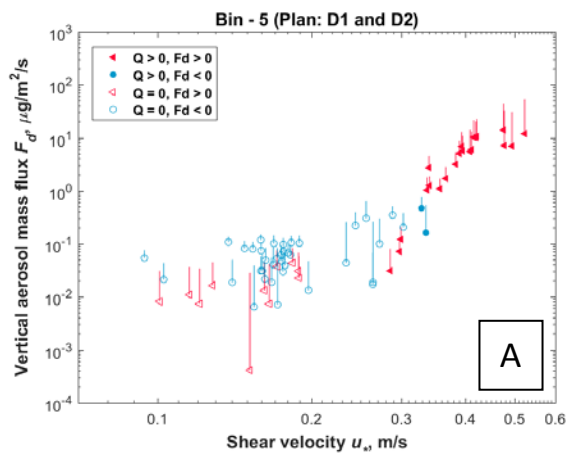
329



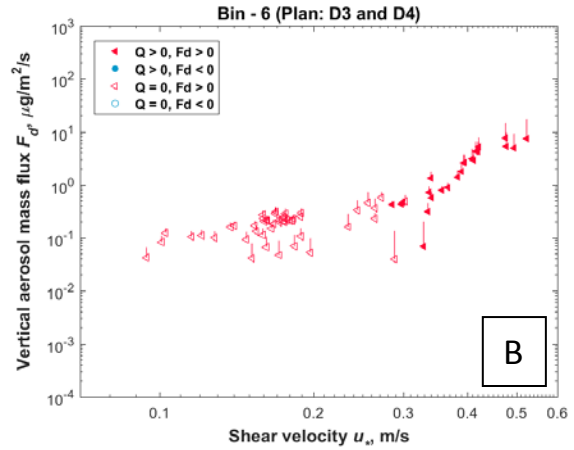
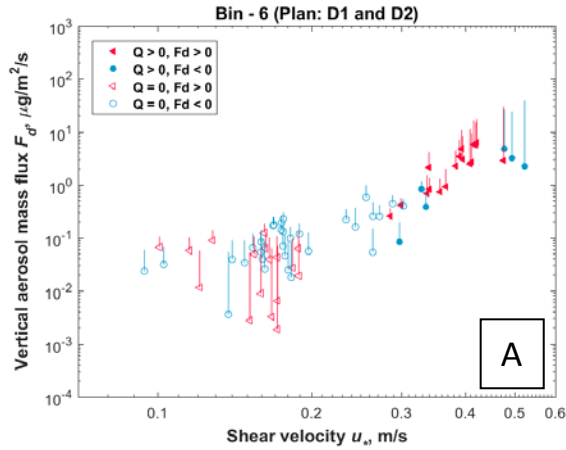
330



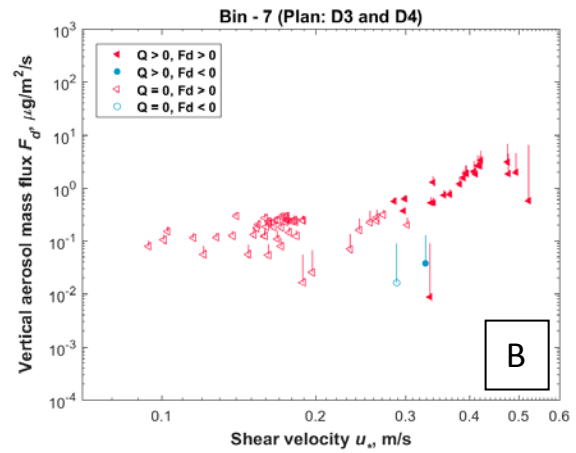
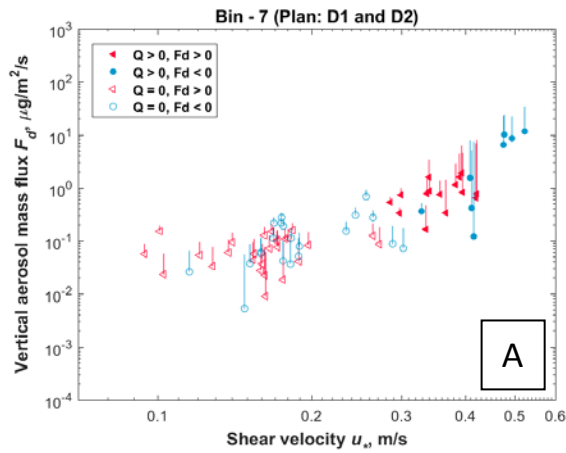
331



332



333

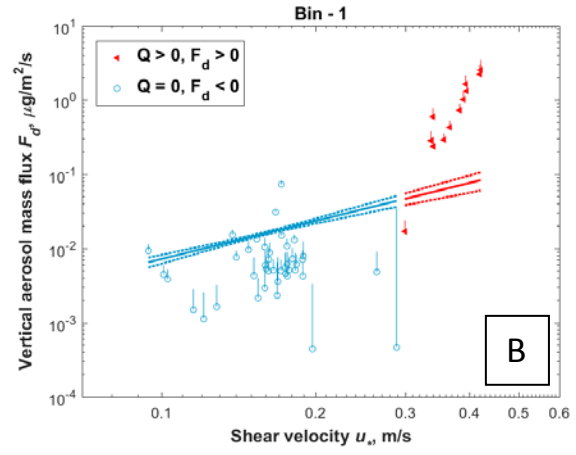
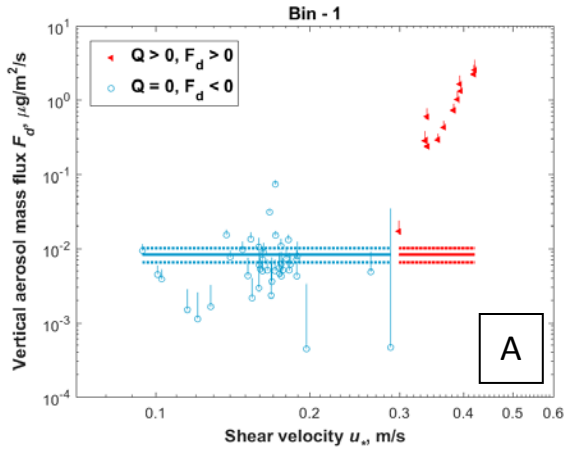


334

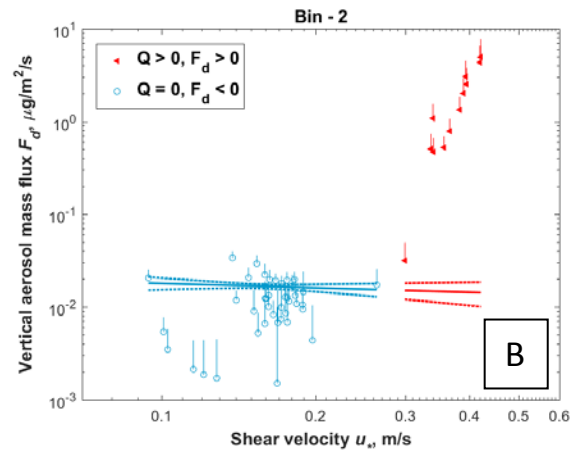
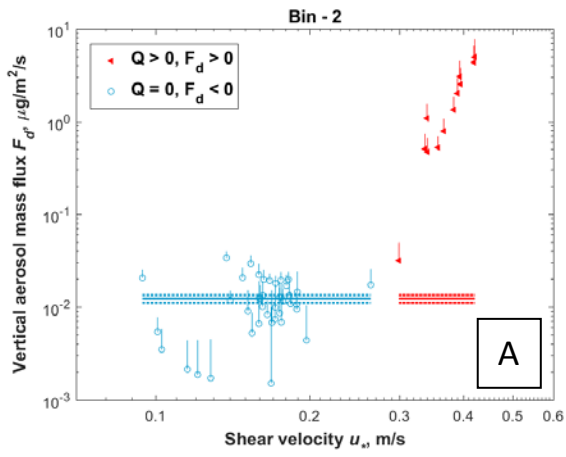
335 Figure S7. Vertical aerosol mass flux as a function of shear velocity calculated by using two plans of optical
 336 particle counters (OPCs)' assemblies during May 29th to June 4th, 2015: (A) using D1 and D2 only, and (B)
 337 using D3 and D4 only. Each of the 14 subplots (2 plans * 7 size bins) includes the four scenarios with active
 338 saltation and positive flux ($Q > 0, F_d > 0$) (red closed left-pointing triangles), $Q > 0$ and $F_d < 0$ (blue
 339 closed circles), $Q = 0$ and $F_d > 0$ (red open left-pointing triangles), and $Q = 0$ and $F_d < 0$ (blue open
 340 circles). Uncertainty range (error bars) from error propagation. Note that we plotted the absolute value of
 341 negative mass flux under logarithmic scale.

342

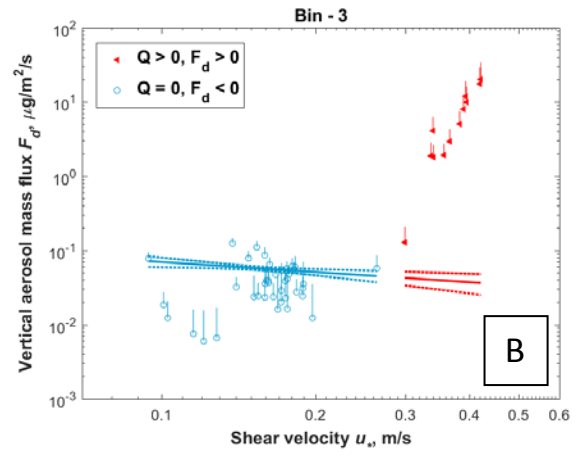
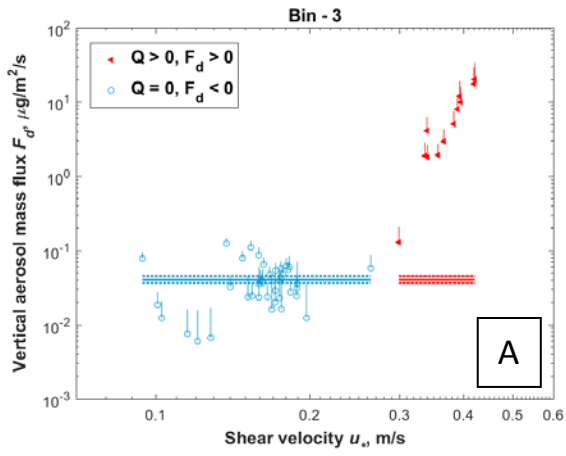
343



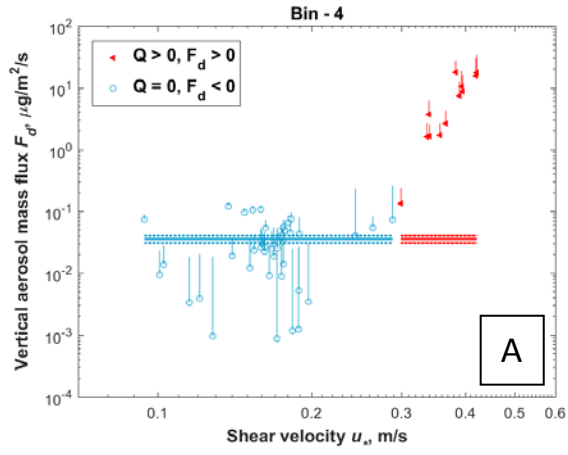
344



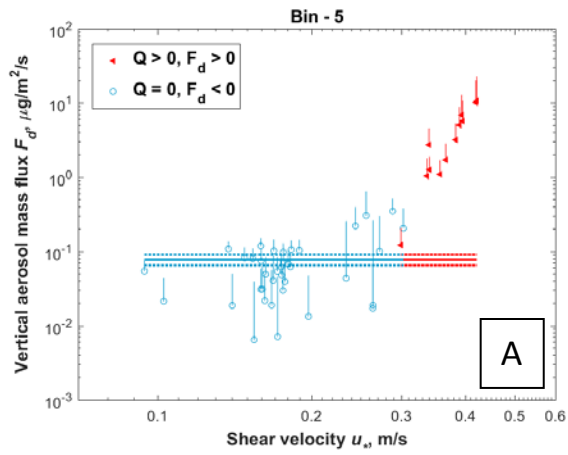
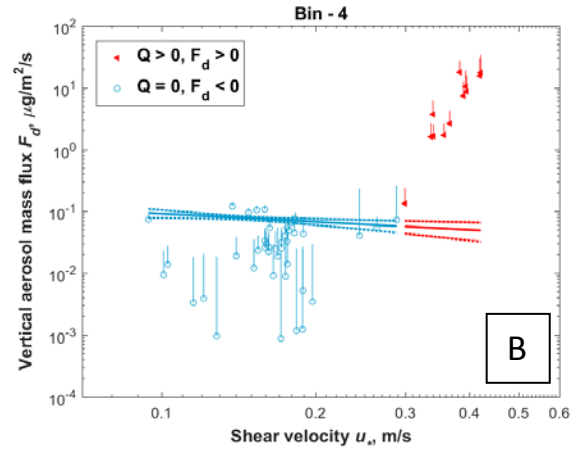
345



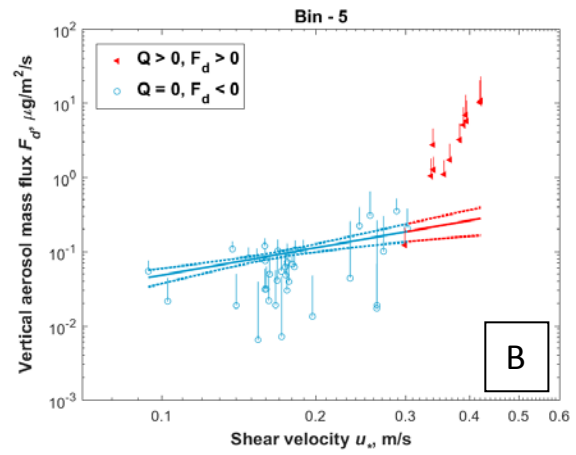
346

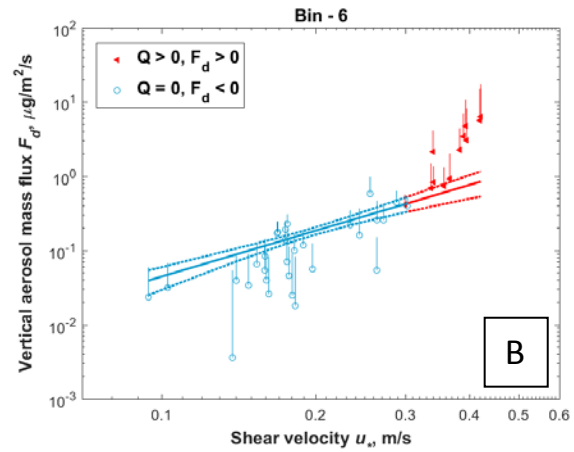
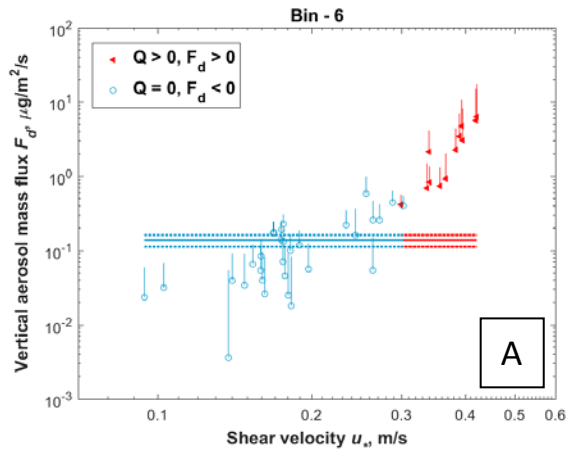


347

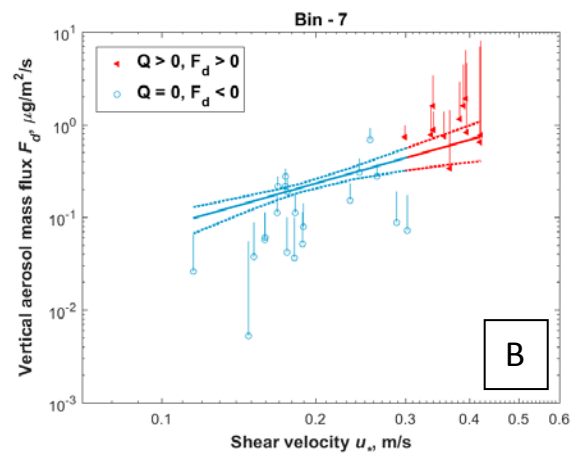
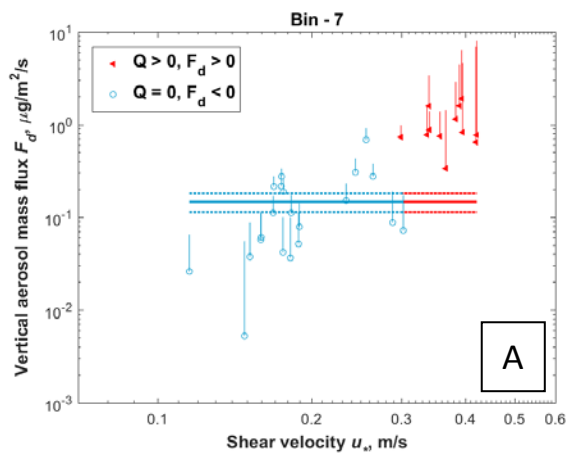


348



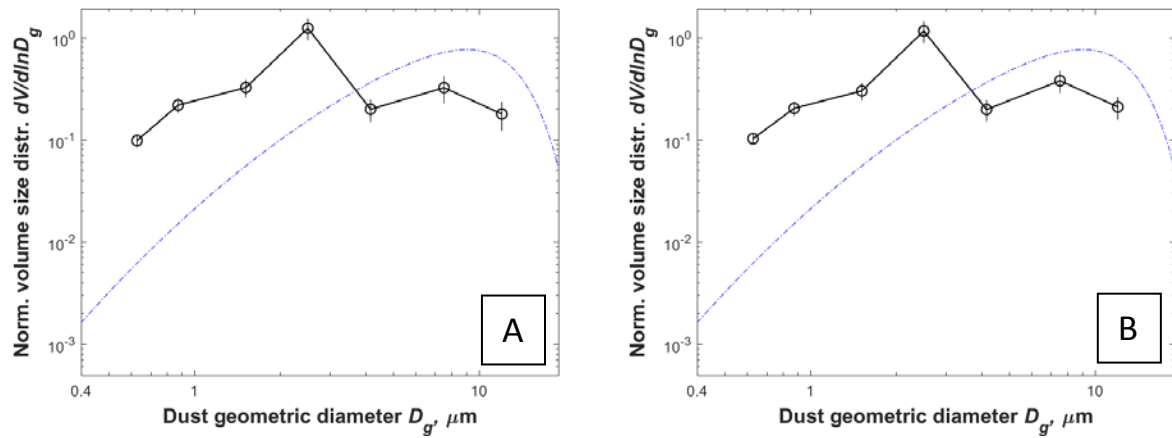


349



350

351 Figure S8. Vertical aerosol mass flux (using D1 and D2 only) as a function of shear velocity during May
 352 29th to June 4th, 2015 using two regression methods: (A) sea-salt deposition flux is invariant with increasing
 353 shear velocity, and (B) sea-salt deposition flux increases non-linearly with shear velocity. Each of the 14
 354 subplots (2 methods * 7 size bins) includes the two scenarios with active saltation and positive flux ($Q >$
 355 $0, F_d > 0$) (red closed left-pointing triangles) and $Q = 0$ and $F_d < 0$ (blue open circles). Uncertainty range
 356 (error bars) from error propagation. Note that we plotted the absolute value of negative mass flux under
 357 logarithmic scale. Results in the main text used the first regression method (A) to remove the deposition of
 358 sea-salt aerosols from our measurements.



359
 360 Figure S9. The normalized volume particle size distribution (PSD) of dust at emission using the mass flux
 361 calculated by assuming that: (A) sea-salt deposition flux is invariant with increasing shear velocity, and that
 362 (B) sea-salt deposition flux increases non-linearly with shear velocity. Also plotted as a reference is the
 363 brittle fragmentation theory (blue dash-dotted lines) on the PSD of emitted dust generated by aggregate
 364 fragmentation (Kok, 2011).