

(Manuscript Ref. NO.: acp-2018-68)

Response to the comments from the Editor and the reviewers

We appreciate the reviewers for their constructive comments on this manuscript. The detailed response to the comments has been listed below.

Overall comments from the Editor:

Thank you for your revision. The reviewers feel that their comments were largely addressed. One however still remains critical about one of the basic assumptions in the manuscript (i.e., absorption is taking place at the dust surface following Henry's law). I agree with this concern and also feel that the answer provided is weak in this regard. I strongly suggest that you attempt to more deeply address this issue in a revision.

Comment from Referee #2:

The authors have thoroughly revised their manuscript taking into account all comments raised by the reviewers. Nevertheless I am still puzzled by the assumption that the gas-dust particles interactions can be description by an absorption processes governed by Henry's law. This would mean that the surface concentration increases linearly with increasing gas phase concentration, which would not fit quite a few observations where a Langmuir-Hinshelwood type behavior has been observed. I acknowledge that the response of the authors that water may form multilayers on top of dust, but this may not confer to it the feature of bulk water. Can you with your assumptions reproduce surface processes as described by a Langmuir-Hinshelwood type equation? If yes, I would then be happy to recommend the publication of your paper.

Response to reviewers:

Our study mainly focused on authentic mineral dust particles (Gobi Desert Dust and Arizona Test dust) under the typical ambient relative humidity (RH) ranging from 20% to 80%. Under the very dry condition, we agree to the reviewers in that the sorption of trace gases and water molecules on mineral dust particles may be governed by a Langmuir-Hinshelwood type adsorption process. However, the observed hygroscopicity suggested that dust particles were coated with the multilayer of water under the RH higher than 20%. Thus, the sorption of trace gases was

approached by the absorption-base model frame (Fig. 2). Several studies also support our assumption showing that the uptake coefficients of trace gases (i.e., SO₂, ozone, and H₂O₂) on dust particles were positively sensitive to RH (0%-90%) (Vlasenko et al., 2006; Yang et al., 2017; Pradhan et al., 2010). Besides, the modification of dust surfaces by forming alkali nitrates can greatly increase hygroscopicity of dust as well as heterogeneous reactions of trace gases.

In order to response to the reviewer, the following sentences have been added to the beginning of Section 3.2 and reads now, “Under the dry condition (RH <20%), the water content on the dust surface is less than a mono-layer (Ibrahim et al., 2018) and it may compete with trace gases for adsorption sites (a Langmuir-like adsorption). However, authentic dust particles which contain inorganic salts and metal oxides, can absorb a significant amount of water molecules under the ambient RH (>20%) *via* a thermodynamic equilibrium process and form a multilayer of water. Besides, the modification of dust surfaces by forming alkali nitrates can greatly increase dust hygroscopicity and furthermore enhance the heterogeneous reactions (Vlasenko et al., 2006; Pradhan et al., 2010; Yang et al., 2017). For example, a higher water content enhances multiphase partitioning of trace gases and the production of oxidized products (HONO, sulfate and nitrate). In the AMAR model, an equation for the dust-phase water content (F_{water} , $\mu\text{g } \mu\text{g}^{-1}$), which is defined as the water mass normalized by the dry dust mass, is mathematically derived for RH between 20% and 80%.”

References

- Ibrahim, S., Romanias, M. N., Alleman, L. Y., Zeineddine, M. N., Angeli, G. K., Trikalitis, P. N., and Thevenet, F.: Water Interaction with Mineral Dust Aerosol: Particle Size and Hygroscopic Properties of Dust, *Acs Earth Space Chem*, 2, 376-386, 10.1021/acsearthspacechem.7b00152, 2018.
- Pradhan, M., Kyriakou, G., Archibald, A. T., Papageorgiou, A. C., Kalberer, M., and Lambert, R. M.: Heterogeneous uptake of gaseous hydrogen peroxide by Gobi and Saharan dust aerosols: a potential missing sink for H₂O₂ in the troposphere, *Atmos Chem Phys*, 10, 7127-7136, 10.5194/acp-10-7127-2010, 2010.
- Vlasenko, A., Sjogren, S., Weingartner, E., Stemmler, K., Gäggeler, H. W., and Ammann, M.: Effect of humidity on nitric acid uptake to mineral dust aerosol particles, *Atmos. Chem. Phys.*, 6, 2147-2160, 10.5194/acp-6-2147-2006, 2006.
- Yang, W. W., Zhang, J. H., Ma, Q. X., Zhao, Y., Liu, Y. C., and He, H.: Heterogeneous Reaction of SO₂ on Manganese Oxides: the Effect of Crystal Structure and Relative Humidity, *Sci Rep-Uk*, 7, ARTN 4550. 10.1038/s41598-017-04551-6, 2017.

Simulation of Heterogeneous Photooxidation of SO₂ and NO_x in the presence of Gobi Desert Dust Particles under Ambient Sunlight

Zechen Yu and Myoseon Jang

P.O.Box116450, Department of Environmental Engineering Sciences, Engineering School of Sustainable
Infrastructure and Environment, University of Florida, Gainesville, FL, USA, 32611

Correspondence to: Myoseon Jang (mjang@ufl.edu)

Abstract

To improve the simulation of the heterogeneous oxidation of SO₂ and NO_x in the presence of authentic mineral dust particles under ambient environmental conditions, the explicit kinetic mechanism was constructed in Atmospheric Mineral Aerosol Reaction (AMAR) model. The formation of sulfate and nitrate was divided into three phases: gas phase, non-dust aqueous phase and dust phase. Specially, AMAR established the mechanistic role of dust chemical characteristics (e.g., photoactivation, hygroscopicity, and buffering capacity) on heterogeneous chemistry. The photo-activation kinetic process of different dust particles was built into the model by measuring the photodegradation rate constant of an impregnated surrogate (malachite green dye) on a dust filter sample (e.g., Arizona Test dust (ATD) and Gobi Desert dust (GDD)) using an online reflective UV-visible spectrometer. The photoactivation parameters were integrated with the heterogeneous chemistry to predict OH radical formation on dust surfaces. A mathematical equation for the hygroscopicity of dust particles was also included in the AMAR model to process the multiphase partitioning of trace gases and in-particle chemistry. The buffering capacity of dust, which is related to the neutralization of dust alkaline carbonates with inorganic acids, was included in the model to dynamically predict the hygroscopicity of aged dust. The AMAR model simulated the formation of sulfate and nitrate using experimental data obtained in the presence of authentic mineral dust under ambient sunlight using a large outdoor smog chamber (UF-APHOR). Overall, both GDD and ATD significantly enhanced the formation of sulfate and nitrate, compared to that in the system without dust particles. However, the influence of GDD on the heterogeneous chemistry was much greater than that of ATD. Based on the model analysis, GDD enhanced the sulfate formation mainly *via* its high photoactivation capability. In the case of NO₂ oxidation, dust-phase nitrate formation is mainly regulated by the buffering capacity of dust. The measured buffering capacity of GDD was two times greater than that of ATD, and consequently, the maximum nitrate concentration with GDD was nearly two times higher than that with ATD. The model also highlights that in urban areas with high NO_x concentrations, hygroscopic nitrate salts quickly form *via* titration of the carbonates in the dust phase, but in the presence of SO₂, the nitrate salts are gradually depleted by sulfate.

1 Introduction

Large quantities of mineral dust particles are frequently ejected into the atmosphere through wind action. Airborne mineral dust is a major contributor to atmospheric particulate matter with an estimated annual emission of 1000-3000 Tg yr⁻¹ (Textor et al., 2006; Tegen and Schepanski, 5 2009). Airborne dust is essential for radiation balance (Sokolik and Toon, 1996; Sokolik et al., 2001; Balkanski et al., 2007), cloud condensation nucleation (Krueger et al., 2003; Liu et al., 2008; Tang et al., 2016), oceanic metal-ion cycles (Jickells et al., 2005; Mahowald et al., 2005; Schulz et al., 2012) and visibility impairment (Kim et al., 2001; Camino et al., 2015).

The surface of mineral dust particles can act as an important sink for atmospheric trace 10 gases, such as O₃, NO_x (e.g., NO and NO₂) and SO₂, and can enhance the production of oxygenated compounds (e.g., nitrate and sulfate) (Underwood et al., 2001; Michel et al., 2002; Usher et al., 2003b; Crowley et al., 2010; George et al., 2015; Tang et al., 2017). For example, 50% to 70% of the annual average total sulfate concentration is estimated to be formed by the heterogeneous oxidation of SO₂ in the vicinity of dust sources (Dentener et al., 1996; Usher et al., 2003a). During 15 a dust event (Beijing, China, on March 28, 2015), Wang et al. (2017) observed that the heterogeneous reactions on dust are the major production mechanisms for nitrate, 19 μg m⁻³, and sulfate, 7 μg m⁻³. Furthermore, airborne dust particles can carry the anthropogenic pollutants to downwind region during the long-range transports (Li and Shao, 2012; Valle-Diaz et al., 2016; Uno et al., 2017).

20 Several recent studies have shown significant increases in sulfate and nitrate concentrations due to the heterogeneous photooxidation of SO₂ and NO_x on mineral dust surfaces. For example, using a flow chamber, Dupart et al. (2014) observed that the NO₂ uptake rate of Arizona Test dust (ATD) particles was 4 times greater under UV-A irradiation than in the dark. A chamber study by Park and Jang (2016) also showed a significantly higher (10 times higher) SO₂ reactive uptake 25 coefficient on ATD under UV light (a mixture of UV-A and UV-B light) than that obtained in the dark. In another chamber study, Park et al. (2017) reported that the increase in the SO₂ kinetic uptake coefficient of Gobi Desert dust (GDD) particles was higher than that observed for the ATD particles. Field observations by Ndour et al. (2009) and Dupart et al. (2012) showed that the uptake coefficients of trace gases (e.g., NO₂ and SO₂) on authentic dust particles increased under sunlight

compare to those in the dark.

Despite numerous studies on the heterogeneous photooxidation of trace gases, the mechanism behind the in-particle chemistry remains largely unknown. One challenge is modeling the photocatalytic process of semi-conductive metal oxides (e.g., TiO_2 and Fe_2O_3) in dust particles. This photocatalytic process results in the formation of electron-hole pairs that can react with a water molecule or absorbed oxygen on the dust surface to form oxidant radicals (e.g., OH radical and HO_2 radical) and oxidize trace gases on dust particles (Linsebigler et al., 1995; Hoffmann et al., 1995; Thompson and Yates, 2006; Cwiertny et al., 2008). Additionally, the hygroscopic property of mineral dust, which is dynamic due to the atmospheric process associated with the dust buffering capacity and inorganic composition, complicates the dust model. For example, Tang et al. (2015) reported decreased hygroscopic properties due to the formation of calcium sulfate *via* the reaction of calcium carbonate with sulfuric acid. Some inorganic salts in dust, such as magnesium sulfate and calcium nitrate, are hydrophilic and can be hydrated in low humidity environments (Liu et al., 2008; Beardsley et al., 2013; Abdelkader et al., 2017). The chemical properties of mineral dust can also be changed by carboxylic acids absorbed on dust particles, which further react with alkaline dust components (Mochizuki et al., 2016). Therefore, deriving a mathematical model to describe the hygroscopicity of dust particles is important for accurately processing both the multiphase partitioning of trace gases and the in-particle chemistry under ambient conditions.

In our recent modeling work (Yu et al., 2017), the heterogeneous oxidation of SO_2 was simulated in the presence of ATD. However, ATD particles have chemical and physical properties that are different from those of ambient mineral dust particles. To simulate the heterogeneous chemistry of trace gases under ambient conditions, a model should include different authentic dusts with various surface areas, hygroscopic properties, photocatalytic capacities, and buffering abilities.

In this study, the Atmospheric Mineral Aerosol Reaction (AMAR) model highlights three aspects to accurately predict the heterogeneous photooxidation of SO_2 and NO_x : (1) the photocatalytic production of OH radicals; (2) the dynamic hygroscopicity of mineral dust; and (3) the buffering capacity determined by the dust compositions. For example, the kinetic mechanisms for the photoactivation processes of different dust particles (ATD and GDD) were established

using the AMAR model based on laboratory data from the photodegradation of an impregnated dye (malachite green) on a dust filter sample. A mathematical model for dust particle hygroscopicity was also integrated into the model based on hygroscopicity data from Fourier transform infrared (FTIR) spectra of fresh and aged particles. The buffering capacity of dust particles was parameterized in the model by measuring the nitrate that formed *via* the photooxidation of NO_x in the presence of dust particles (ATD or GDD) using an indoor chamber with different humidities (20%, 55% and 80%). The resulting AMAR model was then evaluated against chamber data obtained under ambient conditions using a large outdoor smog reactor at the University of Florida Atmospheric Photochemical Outdoor Reactor (UF-APHOR).

10 **2 Experimental section and model description**

2.1 Sample preparation and characterization

The GDD particles were collected from the dust deposition region (Tsogt-Ovoo Soum in the Umnugovi Province, Mongolia) between March and May 2015. The collected sample was sieved to less than 20 μm. The ATD particles are a commercialized dust sample (size range: 0–3 μm) (Power Technology Inc. USA) from Arizona, USA. The particle size distributions of airborne dust particles were measured using both a scanning mobility particle sizer (SMPS; TSI 3080, USA) and an optical particle sizer (OPS; TSI 3330, USA). The measured SMPS and OPS data were merged using the Multi-Instrument Manager (MIM) 2.0 software (TSI, USA). An example of ATD and GDD particle distributions used in this study is shown in Fig. S1. The concentration of geometric surface area (cm² cm⁻³) of airborne dust particles were calculated based on the particle size distribution. The BET surface areas, which were measured using the BET method and a NOVA 2200 instrument, of ATD and GDD were previously reported to be 47.4 and 39.6 m² g⁻¹, respectively (Park et al., 2017). The fraction of elements in the GDD and ATD samples were previously analysed by Park et al. (2017) using an energy dispersive spectroscopy (EDS).

25 2.2 Indoor and outdoor chamber experiments

The indoor and outdoor chamber operations have been previously reported (Yu et al., 2017; Park et al., 2017) (see Sect. S1). To generate the model parameters for the heterogeneous

oxidation of SO₂, preexisting indoor chamber data were employed (Park and Jang, 2016). In this study, nitrate data were added to create the model parameters for NO₂ oxidation. In the presence of different dust particles under various humidity levels (20%, 55% and 80%), NO₂ and SO₂ were photo-oxidized using a 2 m³ indoor Teflon film chamber equipped with 16 UV lamps (wavelength range from 280 nm to 900 nm) (FS40T12/UVB, Solarc Systems Inc., Canada). The details on the experimental conditions for the NO₂ oxidation are listed in Table S1. The resulting AMAR model was tested against the outdoor chamber data produced using the UF-APHOR dual chambers (52 + 52=104 m³) under ambient environmental conditions. The nitrate and sulfate ion concentrations were measured using a particle into-liquid sampler (ADISO 2081, Applikon Inc., Netherlands) coupled with ion chromatography (761 Compact IC, Metrohm Inc., USA) (PILS-IC). The details on the outdoor chamber data are listed in Table 1. The concentrations of NO_x, SO₂ and O₃ were continuously measured using a chemiluminescence NO/NO_x analyzer (Model T201, Teledyne, USA), a fluorescence Total Reduced Sulfur (TRS) analyzer (Model 102E, Teledyne, USA) and a photometric ozone analyzer (Model 400E, Teledyne, USA), respectively.

2.3 Measurement of the dust particle photoactivation parameters

To parameterize the photoactivation capability of dust particles, a dust filter sample impregnated with a dye (malachite green) was photochemically irradiated using a specifically fabricated flow chamber equipped with a UV lamp (11SC-2.12; Pen-Ray., UK) coupled to a cut-off lens ($\leq 280 \pm 5$ nm wavelength, 20CGA-280; Newport, USA) (Fig. S2). The dry dust particles were introduced into the indoor chamber by passing clean air through a nebulizer (Pari LC star, Starnberg, Germany). The dust particles were then collected on a Teflon-coated, glass-fiber filter (Emfab TX40 HI20 WW; Pallflex Corp., Putnam, CT) to obtain 200 μg of dust particles per filter. This filter sample was then impregnated with 4 μg of malachite green dye dissolved in ethanol. Afshar et al. (Afshar et al., 2011) reported that malachite green dye decays in the presence of metal oxides under UV light. The dye-impregnated dust filter sample was placed in a UV flow chamber to activate the heterogeneous photodegradation of the dye on the dust particles. The humidity inside the flow chamber was controlled by manipulating the air flow ($\sim 0.5 \text{ L min}^{-1}$) and passing clean, dry air through a water bubbler. The degradation of the dye impregnated on the dust sample was then measured using a reflective UV-visible spectrophotometer (Jaz Spectrometer; Ocean

Optics Inc., USA). Figure S3 shows an example of measured light absorbance of dye impregnated dust filter before and after irradiation using UV light. The degradation of dye was significant only in the presence of dust particles.

2.4 Hygroscopic properties of dust particles

5 The hygroscopic properties of the fresh and aged dust particles were determined using an FTIR spectrometer (Nicolet Magma 560, Madison, WI, USA) combined with a specifically fabricated optical flow chamber (Zhong and Jang, 2014;Jang et al., 2010;Beardsley et al., 2013;Park et al., 2017) that could control the humidity level in the range from 10% to 80%. The dust particles were impacted onto a silicon FTIR window (13×2 mm; Sigma–Aldrich, St. Louis, 10 MO, USA) and weighed using an analytical balance (MX5; Mettler-Toledo Ltd., England). The FTIR peak at 1650 cm⁻¹ was used to determine the water content of the particles. To calibrate the water content in the dust particles, (NH₄)₂SO₄ particles were used, and the calibration was based on the particle mass and water content estimated using an inorganic thermodynamic model (E-AIM II) (Clegg et al., 1998;Wexler and Clegg, 2002;Clegg and Wexler, 2011).

15 3 Results and discussion

3.1 Description of the AMAR model

The AMAR model was developed to predict the heterogeneous oxidation of SO₂ and/or NO_x in the presence of authentic mineral dust particles. As described in previous work (Yu et al., 2017), the formation of mass concentrations of sulfate ([SO₄²⁻], μg m⁻³) and nitrate ([NO₃⁻], μg m⁻³) is processed in three phases: the gas phase, inorganic salt-seeded aqueous phase and dust phase. 20 The key components of the model consist of multiphase trace gases partitioning and the kinetic mechanisms of the three phases. Under ambient conditions (RH higher than 20%), studies showed that the water content in dust particles ranged 2-4 monolayers based on the BET surface area (Gustafsson et al., 2005;Yu et al., 2017). Therefore, we assume that the gas–dust partitioning of 25 trace gases is governed by the absorption process. The partitioning coefficients of these gases can be calculated using Henry’s Law constant (K_H), and the coefficients are influenced by the dust phase water content. The oxidation of SO₂ and NO_x in the gas phase and inorganic salt-seeded

aqueous phase was simulated using the mechanisms previously reported in the literature (Liang and Jacobson, 1999; Binkowski and Roselle, 2003; Byun and Schere, 2006; Sarwar et al., 2013; Sarwar et al., 2014; Yu et al., 2017). Dust-phase sulfuric acid partially or fully reacts with indigenous alkaline salt or the gaseous ammonia originating from the chamber wall (Li et al., 5 2015; Beardsley and Jang, 2016). The reactive uptake coefficient (γ) of SO₂ or NO₂ on dust in this study is dynamically changing due to changing of sunlight intensity and relative humidity. The calculation of γ is explained in Sect. S2 and an example of γ for GDD particles is shown in Fig. S4.

An overall schematic of the dust-phase chemistry mechanism in the AMAR model is 10 shown in Fig. 1 (also see Table S2). To accurately process the heterogeneous oxidation of SO₂ and NO_x under ambient conditions, we emphasized the three key processes in dust-phase chemistry:

- (1) A mathematical model for dust particle hygroscopicity was derived to dynamically simulate the dust-phase water content as a function of dust aging, e.g., the neutralization of alkaline carbonates and inorganic components containing ammonia, sulfate and nitrate. This 15 hygroscopic model improved the multiphase gas-dust partitioning and in-particle chemistry (Sect. 3.2).
- (2) Kinetic mechanisms to simulate the photoactivation of dust particles and the formation of dust-phase OH radicals were included in the AMAR. Specifically, we standardized the technique to parameterize the photoactivation capability of various dust particles (Sect. 2.3 and Sect. 3.3).
- 20 (3) The neutralization mechanisms for dust particles with inorganic acids were systematically approached using the buffering capacity parameter. This process is linked to the hygroscopicity of dust particles (Sect. 3.4).

3.2 Dust-phase water content

25 Under the dry condition (RH <20%), the water content on the dust surface is less than a mono-layer (Ibrahim et al., 2018) and it may compete with trace gases for adsorption sites (a Langmuir-like adsorption). (Ibrahim et al., 2018) However, authentic dust particles which contain inorganic salts and metal oxides, can absorb a significant amount of water molecules under the ambient RH (>20%) via a thermodynamic equilibrium process and form a multilayer of water

~~— The inorganic salts and metal oxides in dust particles can absorb water via a~~

thermodynamic equilibrium process and form a thin film of water on the dust surface (Vlasenko et al., 2006; Pradhan et al., 2010; Yang et al., 2017). Besides, the modification of dust surfaces by forming alkali nitrates can greatly increase dust hygroscopicity and furthermore enhance the heterogeneous reactions (Vlasenko et al., 2006; Pradhan et al., 2010; Yang et al., 2017). For example, a higher water content enhances multiphase partitioning of trace gases and the production of oxidized products (HONO, sulfate and nitrate). In the AMAR model, an equation for the dust-phase water content (F_{water} , $\mu\text{g } \mu\text{g}^{-1}$), which is defined as the water mass normalized by the dry dust mass, is mathematically derived for RH between 20% and 80%. In general, a higher water content enhances multiphase partitioning of trace gases and the production of oxidized products (HONO, sulfate and nitrate). In the AMAR model, an equation for the dust-phase water content (F_{water} , $\mu\text{g } \mu\text{g}^{-1}$), which is defined as the water mass normalized by the dry dust mass, is mathematically derived. F_{water} is estimated by an additive function with three parts:

$$F_{water} = a(e^{b \cdot RH} - 1) + c \cdot e^{d \cdot RH} \frac{[NO_3^-]}{[Dust]} + \frac{[Water]_{SO_4^{2-} - NH_4^+ - H_2O}}{[Dust]} \quad (1)$$

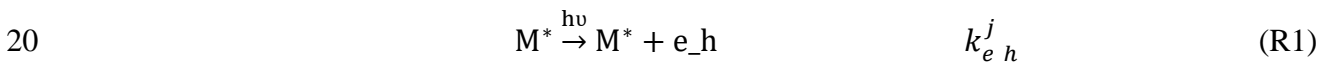
where RH represents the relative humidity and ranges from 0.2 to 0.8. The first term, $a(e^{b \cdot RH} - 1)$, in Eq. (1) is associated with the water content of fresh dust particles. The 2nd term, $c \cdot e^{d \cdot RH} \frac{[NO_3^-]}{[Dust]}$, represents the hygroscopicity of the hydrophilic nitrate salts that are formed via titration of the dust constituents (e.g., alkaline carbonates and some metal oxides). The 3rd term, $\frac{[Water]_{SO_4^{2-} - NH_4^+ - H_2O}}{[Dust]}$, originates from the ammonium sulfate system and is estimated via the inorganic thermodynamic model E-AIM II (Clegg et al., 1998; Wexler and Clegg, 2002; Clegg and Wexler, 2011). Coefficients a (0.03 ± 0.01), b (3.6 ± 0.5), c (1.4 ± 0.4) and d (4.0 ± 0.4) are dimensionless and they were determined using FTIR data (Fig. 2).

To determine the coefficients a and b, the hygroscopicity of fresh ATD particles or fresh GDD particles was measured using an FTIR spectrometer for RH levels from 10% to 80%. Similarly, the coefficients c and d were obtained from the FTIR spectra of aged dust particles, e.g., NO₂ photooxidation in the presence of ATD particles or GDD particles. The nitrate concentrations ($\mu\text{g } \mu\text{g}^{-1}$ in dust mass) were measured using PILS-IC and were 0.001 (approximately negligible) for fresh ATD and 0.011 for aged ATD. The nitrate concentrations were 0.007 for fresh authentic GDD and 0.02 for aged GDD. Figure 2 shows the F_{water} values for ATD and GDD particles with

and without aging. For both the fresh and aged dust particles, the F_{water} value gradually increases in the dry region (RH < 40%) but rapidly increases for RH values greater than 40%. The F_{water} value of fresh GDD is higher than that of ATD for the entire RH range due to the presence of more hydrophilic nitrate salts. Assuming that the F_{water} value from the 2nd term has a linear relationship with the nitrate content, the F_{water} value associated with nitrate salts can be estimated. Figure 2(b) shows that when the nitrate-associated F_{water} is excluded, the F_{water} value of fresh GDD (e.g., the hygroscopicity solely originating from dust constituents other than nitrates) is similar to that of ATD. The difference in model parameters for hygroscopicity between ATD and GDD is insignificant. For dust aged *via* the photooxidation of NO₂, the GDD particles are more hygroscopic than the ATD particles. Overall, clear phase transitions and obvious differences between the hydration and dehydration processes were not observed for either type of dust particle. This trend suggests that the hygroscopicity of dust particles is caused by a variety of chemical species.

3.3 ATD and GDD photoactivation parameters

Mineral dust plays a key mechanistic role as a photocatalyst to accelerate the oxidation of trace gases in the dust phase. The photoactivation of semiconducting metal oxides (M*) in dust particles can yield an electron-hole pair ($e^-_{cb}-h^+_{vb}$) that further reacts with water or oxygen molecules to form oxidizing radicals, such as OH radicals (Linsebigler et al., 1995; Hoffmann et al., 1995; Thompson and Yates, 2006; Cwiertny et al., 2008; Yu et al., 2017).



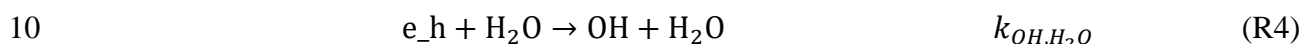
where e_h is an $e^-_{cb}-h^+_{vb}$ pair and $k_{e_h}^j$ is the operational photoactivation rate constant of dust particles. The production rate of the $e^-_{cb}-h^+_{vb}$ pair is described as

$$\frac{d[e_h]_{dust}}{dt} = k_{e_h}^j [M^*]_{dust} \quad (2)$$

where $[M^*]_{dust}$ is the concentration (molecules cm⁻²) of M* on the dust surface (calculate based on geometric surface area). In our recent study (Yu et al., 2017), $k_{e_h}^j$ was linked to the wave-dependent, mass-absorbance cross section and quantum yield of a given dust particle (ATD) (Fig. S5). However, the type and quantity of conductive constituents in authentic dust particles vary. Hence, to extend the model to ambient conditions, the photoactivation of different dust particles

and their kinetic mechanisms must be scaled.

In this study, we determined the relative photoactivation rate constant for different dust particles using colorimetry integrated with a fabricated photochemical flow reactor (also see Sect. 2.3). The impregnated dye (malachite green) on the dust surface was photodegraded by the oxidants created by the dust particles. The relative degradation rate constant of the dyed filter was measured using an online reflective UV-visible spectrometer to scale the photoactivation of the dust. The kinetic mechanisms for the reactions of the dye with radicals are expressed as follows



where k_{recom} is the rate constant of the recombination reaction of an electron with a hole. The concentration of the dye on the dust surface was assumed to be significantly higher than that of the surface OH radical and much lower than that of the water content or oxygen molecules on dust. The concentration unit (molecules cm^{-2}) of the chemical species in R1-R5 was multiplied by the surface area concentration of the dust particles (A_{dust} , $\text{cm}^2 \text{cm}^{-3}$) to convert the value to the concentration unit in air (molecules cm^{-3}). By combining R1-R5, the kinetic reaction rates for the $e^-_{cb}-h^+_{vb}$ pairs, OH radicals and dye can be written as.

$$\frac{d[e_h]_{air}}{A_{dust}dt} = k_{e_h}^j \frac{[M^*]_{air}}{A_{dust}} - k_{recom} \frac{[e_h]_{air}}{A_{dust}} - k_{OH,O_2} \frac{[e_h]_{air}[O_2]_{air}}{A_{dust}^2} - k_{OH,H_2O} \frac{[e_h]_{air}[H_2O]_{air}}{A_{dust}^2} \quad (3)$$

$$20 \quad \frac{d[OH]_{air}}{A_{dust}dt} = k_{OH,O_2} \frac{[e_h]_{air}[O_2]_{air}}{A_{dust}^2} + k_{OH,H_2O} \frac{[e_h]_{air}[H_2O]_{air}}{A_{dust}^2} - k_{dye} \frac{[OH]_{air}[dye]_{air}}{A_{dust}^2} \quad (4)$$

$$\frac{d[dye]_{air}}{A_{dust}dt} = -k_{dye} \frac{[dye]_{air}[OH]_{air}}{A_{dust}^2} \quad (5)$$

The concentration of chemicals with the subscript “air” is the concentration in air (molecules cm^{-3}). Under the assumption of a steady state for the net reaction rate of an $e^-_{cb}-h^+_{vb}$ pair and OH radical, the dye consumption rate can be written as

$$25 \quad \frac{d[dye]_{air}}{dt} = - \frac{k_{e_h}^j [M^*]_{air}}{\frac{k_{recom}}{k_{OH,O_2}[O_2]_{dust} + k_{OH,H_2O}[H_2O]_{dust}} + 1} \quad (6)$$

where k_{recom} is much larger than $k_{OH,O_2}[O_2]_{dust}$ or $k_{OH,H_2O}[H_2O]_{dust}$ in Eq. (6). A previous study by Khorasani et al. (2014) also reported a recombination rate ($\sim 10^4 \text{ s}^{-1}$) for an $e^-_{cb}-h^+_{vb}$ pair on

silicon that is much faster than the rate observed for typical in-particle reactions. Therefore, the term $\left(\frac{k_{recom}}{k_{OH,O_2}[O_2]_{dust}+k_{OH,H_2O}[H_2O]_{dust}}\right)$ in Eq. (6) is much larger than 1. $[O_2]_{dust}$ is calculated through the partitioning process as follows

$$[O_2]_{dust} = K_p [O_2]_{gas} [H_2O]_{dust} \quad (7)$$

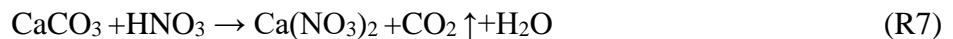
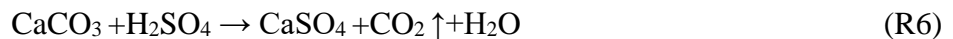
5 where $[O_2]_{gas}$ is the concentration of oxygen in the air and K_p is the partitioning coefficient for O_2 on the dust-phase water layer. By applying Eq. (7) to (6), the analytical solution for Eq. (6) can be written as

$$\Delta[dye]_{dust} = -k_{e,h}^j [M^*]_{dust} [H_2O]_{dust} \left(\frac{k_{OH,O_2} K_p [O_2]_{gas} + k_{OH,H_2O}}{k_{recom}}\right) t \quad (8)$$

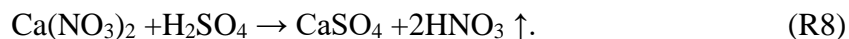
As shown in Eq. (8), the dye decomposition on the particle surface is proportional to $[M^*]$, which
 10 changes based on the dust type, and the dust-phase water concentration, which can be estimated using F_{water} and changes with the dust composition. Figure 3 shows the dye degradation rate in the presence of ATD or GDD particles, and the rate was measured using a UV flow chamber (Sect. 2.3). The $[M^*]$ value, which leverages the photoactivation ability of dust particles, is included in Fig. (3). The $[M^*]$ value of GDD is 2.5 ± 0.7 times higher than that of ATD. This difference in
 15 dust's photoactivation ability can be explained by the dissimilarity in their elemental compositions. As seen in the previous study by (Park et al., 2017), the elemental fraction of conductive metals such as iron and tritium appeared to be higher with the GDD of this study than reference ATD. The correlation between the metal compositions and photoactivation ability of dust particles needs to be explored in future. Additionally, the estimated photoactivation parameters of both GDD and
 20 ATD particles increase with increasing humidity suggesting the importance of the role of water molecules to heterogeneous oxidation reactions.

3.4 Impact of the dust buffering capacity

The buffering capacity is determined by the neutralization of the dust-phase constituents (e.g., alkaline carbonates and some metal oxides) with inorganic acids. For example, alkaline
 25 carbonates in dust particles can react with nitric acid or sulfuric acid to form alkaline salts.



In contrast to nitrate, sulfate can accumulate at levels beyond the neutralization capacity of dust because sulfuric acid is not volatile in ambient humidity levels. Furthermore, sulfuric acid can deplete the nitrate salts that build up in the dust phase *via* the following reaction.



5 The buffering capacity determines the maximum nitrate concentration that can build up on dust particles. Nitrate ions are hydrophilic and significantly influence the hygroscopicity of dust particles (F_{water} in Sect. 3.2). The buffering capacity was incorporated into the kinetic mechanisms in the AMAR model to dynamically modulate the F_{water} value.

The buffering capacities of two different mineral dusts (ATD and GDD) were semi-empirically determined by fitting the nitrate prediction to the experimental data shown in Fig. S6(experimental conditions in Table S1) using the kinetic mechanisms (R7 and gas-particle nitric acid partitioning) in the AMAR model. The buffering capacity was determined using the maximum nitrate salt mass normalized by the dust mass (Sect. 2.1). The measured buffering capacities of ATD and GDD are $0.011 \mu\text{g } \mu\text{g}^{-1}$ and $0.020 \mu\text{g } \mu\text{g}^{-1}$, respectively. The difference in buffering capacity between GDD and ATD originates from the content of alkaline carbonates and partially metal oxides.- The element analysis measured by (Park et al., 2017) showed that GDD contained the greater amount of alkaline metals (e.g., K, Ca, Na and Mg) and transition metals (e.g., Fe and Ti) than ATD. However, the reaction generally occurs on the surface of dust rather than the whole body of dust due to its solidity and tortuosity. Thus, the actual buffering capacity of dust is much smaller than the total amount of alkaline carbonates and metal oxides in bulk dust.

3.5 Simulation of outdoor chamber data using the AMAR model

The resulting AMAR model was tested against the outdoor chamber data obtained from simulating the oxidation of NO_x (Fig. 4(a) and 4(a)) or SO_2/NO_x (Fig. 4(c) and 4(d)) in the presence of mineral dust particles under ambient sunlight for 10 hours. As shown in Fig. 4(a) and 25 4(b), nitrate rapidly formed in the morning, and the model well modeled the chamber data. Additionally, the nitrate mass normalized by the dust mass was higher for GDD than ATD. In addition, nitrate depletion was observed (Fig. 4(a) and 4(b)) even in the absence of SO_2 . The nitrate depletion in the chamber data is possibly due to the nitrate salts reacting with the carboxylic acids present in the chamber air, but the current model cannot model this reaction. As shown in Fig. 4(c)

and 4(d), the model well predicts the sulfate and nitrate concentrations produced from SO₂ oxidation at two different NO_x levels in the presence of GDD particles. The oxidation of SO₂ was suppressed when the NO_x concentration was high because SO₂ competes with NO₂ to react with the OH radicals that form on dust surfaces. In the presence of SO₂, the model reasonably predicts the nitrate profile and shows that the nitrate quickly builds up in the morning and is significantly depleted by the formed sulfate. The of SO₂, NO_x, ozone and dust particle concentrations are simulated in Fig. S7.

Figure 4 illustrates the predicted F_{water} values with aging. The F_{water} value is mainly influence by the humidity, which is high in the morning and gradually decreases as the temperature increases. However, the value is also modulated by the mineral dust particle aging process. For example, although the humidity level decreases between 8 AM and 10 AM, the F_{water} value noticeably increases and coincides with the hygroscopic nitrate concentration time profile. The F_{water} value is significantly lower in the presence of SO₂ (Fig. 4(d)) than its absence (Fig. 4(b)) because the sulfate salts on dust particles and sulfates with ammonium ions (e.g., more titrated than ammonium hydrogen sulfate) are less hygroscopic than nitrate salts.

3.6 Model sensitivity

The sensitivity of the model predictions for nitrate (Fig. (5)) and sulfate (Fig. (6)) to the major input variables (e.g., relative humidity, temperature, sunlight intensity, dust mass concentration and NO_x concentration) was evaluated. The sensitivity test was mainly performed for GDD particles (100 μg m⁻³) under the environmental conditions at Gainesville, Florida on 23 November 2017. The nitrate and sulfate mass concentrations in Fig. 5 and 6, respectively, are normalized with the dust mass.

As shown in Fig. 5(a) and 6(a), the formation of both nitrate and sulfate is significantly sensitive to the RH level, but the reasons for this sensitivity are different. There was a sudden increase in the nitrate concentration between a low RH (20% and 55%) and a high RH (80%), imitating the F_{water} trend. In addition to the nitrate salt formation, which is influenced by the buffering capacity, the partitioning of hydrophilic nitric acid into the water layer increases at a higher RH. Unlike nitrate, the sulfate concentration gradually increases as the RH increases. The F_{water} value of the sulfate salts (Fig. 6(a)) is relatively smaller than that of the nitrate salts (Fig.

5(a)). Additionally, nitrate formation is more sensitive to temperature than sulfate formation due to the nitric acid partitioning process. For the different dust types (ATD vs. GDD), the formation of nitrate (Fig. 5(c)) and sulfate (Fig. 6(c)) is higher with GDD. As discussed in Sect. 3.4, the maximum amount of nitrate salts in the dust phase is determined by the buffering capacity of the dust particles. The buffering capacity of GDD is two times higher than that of ATD (e.g., $0.011 \mu\text{g } \mu\text{g}^{-1}$ for ATD and $0.020 \mu\text{g } \mu\text{g}^{-1}$ for GDD), and thus, the nitrate concentration in the GDD system is nearly two times higher than that in the ATD system. Another reason for the high sulfate formation in the presence of GDD is the photoactivation ability of GDD. An in-depth explanation will be presented in Sect. 4. The sunlight intensity has more of an impact on sulfate formation than nitrate formation, as seen in Fig. 5(d) and Fig. 6(d). Although nitrate formation is accelerated by strong sunlight, the nitrate production under the different sunlight intensities is governed by the buffering capacity of a given dust type (e.g., GDD).

Figure 6(e) shows the sulfate formation sensitivity to three levels of NO_x (2, 20 and 40 ppb NO_x) under ambient conditions (e.g., sunlight, temperature and humidity). In the presence of NO_x (40 ppb), the sulfate formation sensitivity to three different RH levels (20, 55, and 80%) was tested, as shown in Fig. 6(f). In general, sulfate is suppressed by increasing NO_x concentrations (Fig. 6(e)). Similar to the effects of humidity on nitrate production at a low NO_x level (Fig. 6(a)), the nitrate formation with a higher NO_x concentration (40 ppb) is also enhanced by a higher RH level, as seen in Fig. 6(f). Additionally, Fig. 6(g) shows how the total sulfate can be attributed to sulfate originating from the reactions in the different phases: (1) the sulfate from the gas phase and inorganic salt-seeded aqueous phase and (2) the sulfate from the dust phase. Dust-phase sulfate formation is suppressed by NO_x due to competition between the absorbed SO_2 and NO_2 for surface OH radicals, while sulfate formation in the inorganic salt-seeded aqueous phase is promoted by NO_x . When the %RH increases from 20 to 80, the heterogeneous reaction is significantly promoted due to the large F_{water} value that enhances both the partitioning process and the production of OH radicals on dust surfaces.

4 Model uncertainties

To characterize the impact of dust characteristics on sulfate formation, the heterogeneous oxidation of SO₂ in the presence of five different dust types, including ATD, GDD, and three artificially formulated dusts (Dust I, II and III), was compared. As shown in Fig. 7, the three characteristic parameters of the dust particles, including the photoactivation capability ($[M^*]_{\text{dust}}$ in Eq. (2)), buffering capacity (Sect. 3.4), and hygroscopicity (F_{water} in Sect. 3.2), were scaled relative to the ATD particles. The relative values of the three parameters for GDD were obtained using laboratory data. The simulation with three artificially formulated samples, Dust I, II, and III, was used to analyze why GDD particles have a larger influence on sulfate formation than ATD and which dust characteristic parameters are the most important for sulfate formation. Figure 7 illustrates that the photoactivation ability of dust ($[M^*]_{\text{dust}}$) is the most important among the three parameters. For example, the sulfate formation noticeably increased between ATD and Dust I. When the three characteristic dust parameters are determined by laboratory studies, in the future, the model can simulate the impact of authentic dust particles on sulfate formation.

Figure S8 also shows the uncertainty in the sulfate and nitrate predictions in the presence of GDD using the AMAR model based on three major dust characteristic parameters (e.g., F_{water} , buffering capacity and photoactivation capability). Assuming that the sulfuric acid beyond the buffering capacity of GDD is treated by the NH₄⁺-SO₄²⁻-H₂O system, we estimated the F_{water} value using an inorganic thermodynamic model with a large uncertainty (E-AIM II) (Clegg et al., 1998; Wexler and Clegg, 2002). In the model simulation, the $\pm 10\%$ uncertainty in F_{water} result in $\pm 7.8\%$ variation in the sulfate concentration and -0.9% to 1.2% variation in the nitrate concentration. As shown in Fig. 3, the uncertainty in the photoactivation parameters of dust particles varies with the RH; e.g., the uncertainty is higher at higher RH levels. The probable uncertainty for the photoactivation of GDD particles at a high %RH (80%) is $\pm 50\%$ and results in -47.7% to 55.7% variation in the sulfate concentrations and -1.0% to 1.9% in the nitrate concentration. The uncertainty in the buffering capacity ($\pm 10\%$) is associated with using ion chromatography to measure the ion concentrations and yields -0.7% to 0.8% variation in the sulfate concentration and -7.6% to 9.4% in the nitrate concentration.

5 Atmospheric implications

Dust storms originating from the Gobi Desert often outbreak during the spring season and influence the air quality over polluted urbans or industrial areas in East Asia (Hsu et al., 2010; Li et al., 2012). In a typical, polluted urban environment; e.g., where NO_x and SO_2 levels are high (40 ppb of NO_x , 5 ppb of SO_2 and $200 \mu\text{g m}^{-3}$ of GDD), the AMAR model shows authentic dust particles are quickly saturated with nitrate and sulfate (concentrations higher than the buffering capacity of GDD), as shown in Fig. S9(a). Under the high NO_x conditions in most urban areas, the heterogeneously formed nitrate on the dust particles modulates the dust hygroscopicity, which is generally higher than that of unaged dust particles. Under high SO_2 concentrations (e.g., 20 ppb of SO_2 , Fig. S9(b)), the dust-phase sulfate depletes nitrate, as discussed in Sect. 3.4. Therefore, SO_2 and NO_x rapidly convert into the nitrate or sulfate during dust break episodes. It is known that Inorganic acids can corrode metal oxides, but they first react with alkaline carbonates on dust. Additionally, the excess amount of sulfuric acid beyond the dust buffering capacity can be titrated by ammonia, which is ubiquitous in ambient air. Thus, the acidity of dust particles may not be high enough to damage the photocatalytic ability of mineral dust particles under ambient conditions. Based on our simulation (Fig. S9), it takes 5-6 days under the ambient conditions to produce the similar amount of sulfate observed in chamber studies (Fig. 4) and it is even longer than the reported average lifetime of airborne dust particles (~ 4.3 days) (Scheuvens and Kandler, 2014; Zender et al., 2004). Therefore, most dust particles possibly would settle down before they are significantly corroded by sulfuric acid coating.

Under ambient conditions, the photooxidation of hydrocarbons in the presence of NO_x is indispensable for the formation of ozone. In the model, the absorbed ozone on dust surfaces positively modulates the formation of sulfate and nitrate *via* either the autoxidation mechanism or the production of OH radicals (Yu et al., 2017). Although NO_2 generally suppresses the formation of sulfate, its influence on heterogeneous chemistry of SO_2 is compounded with ozone in ambient air. For example, heterogeneous chemistry of ozone becomes important in nighttime, particularly when humidity is high, and promotes SO_2 oxidation. Additionally, some organic compounds can sink onto dust surfaces *via* a partitioning process and complicate the heterogeneous chemistry in the model. For example, the organic carboxylic acids on dust surfaces can react with alkaline carbonates to form alkaline carboxylates. Beardsley et al. (2013) reported that anions in inorganic

aerosols, such as NO_3^- , can be depleted by the formation of carboxylic acids, and the subsequent change aerosol hygroscopic properties. Semivolatile organic compounds compete with the absorbed SO_2 and NO_2 for the consumption of OH radicals. Therefore, the model requires further in-depth dust chemistry of organic compounds in the future to accurately predict sulfate and nitrate formation in ambient environments.

Acknowledgments

This research was supported by the award from the Ministry of Science and ICT, the Ministry of Environment, the Ministry of Health and Welfare (2017M3D8A1090654), the award from the National Institute of Metrological Sciences (KMA2018-00512), and the Scholarship from the Fulbright Scholar (from USA to Mongolia).

References

- Abdelkader, M., Metzger, S., Steil, B., Klingmuller, K., Tost, H., Pozzer, A., Stenchikov, G.,
Barrie, L., and Lelieveld, J.: Sensitivity of transatlantic dust transport to chemical aging
and related atmospheric processes, *Atmos Chem Phys*, 17, 3799-3821, 10.5194/acp-17-
3799-2017, 2017.
- 5 Afshar, S., Jahromi, H. S., Jafari, N., Ahmadi, Z., and Hakamizadeh, M.: Degradation of malachite
green oxalate by UV and visible lights irradiation using Pt/TiO₂/SiO₂ nanophotocatalyst,
Sci Iran, 18, 772-779, 10.1016/j.scient.2011.06.007, 2011.
- 10 Balkanski, Y., Schulz, M., Claquin, T., and Guibert, S.: Reevaluation of Mineral aerosol radiative
forcings suggests a better agreement with satellite and AERONET data, *Atmos Chem Phys*,
7, 81-95, 2007.
- Beardsley, R., Jang, M., Ori, B., Im, Y., Delcomyn, C. A., and Witherspoon, N.: Role of sea salt
aerosols in the formation of aromatic secondary organic aerosol: yields and hygroscopic
properties, *Environ Chem*, 10, 167-177, 10.1071/En13016, 2013.
- 15 Beardsley, R. L., and Jang, M.: Simulating the SOA formation of isoprene from partitioning and
aerosol phase reactions in the presence of inorganics, *Atmos Chem Phys*, 16, 5993-6009,
10.5194/acp-16-5993-2016, 2016.
- Binkowski, F. S., and Roselle, S. J.: Models-3 Community Multiscale Air Quality (CMAQ) model
aerosol component 1. Model description, *Journal of geophysical research: Atmospheres*,
20 108, 2003.
- Byun, D., and Schere, K. L.: Review of the governing equations, computational algorithms, and
other components of the models-3 Community Multiscale Air Quality (CMAQ) modeling
system, *Appl Mech Rev*, 59, 51-77, 10.1115/1.2128636, 2006.
- 25 Camino, C., Cuevas, E., Basart, S., Alonso-Pérez, S., Baldasano, J., Terradellas, E., Marticorena,
B., Rodríguez, S., and Berjón, A.: An empirical equation to estimate mineral dust
concentrations from visibility observations in Northern Africa, *Aeolian Research*, 16, 55-
68, 2015.
- Clegg, S., and Wexler, A. S.: Densities and Apparent Molar Volumes of Atmospherically
Important Electrolyte Solutions. 2. The Systems H⁺-HSO₄⁻-SO₄²⁻-H₂O from 0 to 3 mol
30 kg⁻¹ as a Function of Temperature and H⁺-NH₄⁺-HSO₄⁻-SO₄²⁻-H₂O from 0 to 6 mol
kg⁻¹ at 25° C Using a Pitzer Ion Interaction Model, and NH₄HSO₄-H₂O and (NH₄)₃H
(SO₄)₂-H₂O over the Entire Concentration Range, *J Phys Chem A*, 115, 3461-3474, 2011.
- Clegg, S. L., Brimblecombe, P., and Wexler, A. S.: Thermodynamic model of the system H⁺-
NH₄⁺-SO₄²⁻-NO₃⁻-H₂O at tropospheric temperatures, *J Phys Chem A*, 102, 2137-2154, DOI
35 10.1021/jp973042r, 1998.
- Crowley, J. N., Ammann, M., Cox, R. A., Hynes, R. G., Jenkin, M. E., Mellouki, A., Rossi, M. J.,
Troe, J., and Wallington, T. J.: Evaluated kinetic and photochemical data for atmospheric
chemistry: Volume V - heterogeneous reactions on solid substrates, *Atmos Chem Phys*, 10,
9059-9223, 10.5194/acp-10-9059-2010, 2010.
- 40 Cwiertny, D. M., Young, M. A., and Grassian, V. H.: Chemistry and photochemistry of mineral
dust aerosol, *Annu. Rev. Phys. Chem.*, 59, 27-51,
10.1146/annurev.physchem.59.032607.093630, 2008.

- Dentener, F. J., Carmichael, G. R., Zhang, Y., Lelieveld, J., and Crutzen, P. J.: Role of mineral aerosol as a reactive surface in the global troposphere, *J Geophys Res-Atmos*, 101, 22869-22889, Doi 10.1029/96jd01818, 1996.
- 5 Dupart, Y., King, S. M., Nekat, B., Nowak, A., Wiedensohler, A., Herrmann, H., David, G., Thomas, B., Miffre, A., Rairoux, P., D'Anna, B., and George, C.: Mineral dust photochemistry induces nucleation events in the presence of SO₂, *P Natl Acad Sci USA*, 109, 20842-20847, 10.1073/pnas.1212297109, 2012.
- 10 Dupart, Y., Fine, L., D'Anna, B., and George, C.: Heterogeneous uptake of NO₂ on Arizona Test Dust under UV-A irradiation: An aerosol flow tube study, *Aeolian Research*, 15, 45-51, 10.1016/j.aeolia.2013.10.001, 2014.
- George, C., Ammann, M., D'Anna, B., Donaldson, D. J., and Nizkorodov, S. A.: Heterogeneous Photochemistry in the Atmosphere, *Chem Rev*, 115, 4218-4258, 10.1021/cr500648z, 2015.
- 15 Gustafsson, R. J., Orlov, A., Badger, C. L., Griffiths, P. T., Cox, R. A., and Lambert, R. M.: A comprehensive evaluation of water uptake on atmospherically relevant mineral surfaces: DRIFT spectroscopy, thermogravimetric analysis and aerosol growth measurements, *Atmos Chem Phys*, 5, 3415-3421, DOI 10.5194/acp-5-3415-2005, 2005.
- Hoffmann, M. R., Martin, S. T., Choi, W. Y., and Bahnemann, D. W.: Environmental Applications of Semiconductor Photocatalysis, *Chem Rev*, 95, 69-96, DOI 10.1021/cr00033a004, 1995.
- 20 Hsu, S.-C., Wong, G. T., Gong, G.-C., Shiah, F.-K., Huang, Y.-T., Kao, S.-J., Tsai, F., Lung, S.-C. C., Lin, F.-J., and Lin, I.: Sources, solubility, and dry deposition of aerosol trace elements over the East China Sea, *Marine Chemistry*, 120, 116-127, 2010.
- Ibrahim, S., Romanias, M. N., Alleman, L. Y., Zeineddine, M. N., Angeli, G. K., Trikalitis, P. N., and Thevenet, F.: Water Interaction with Mineral Dust Aerosol: Particle Size and Hygroscopic Properties of Dust, *Acs Earth Space Chem*, 2, 376-386, 10.1021/acsearthspacechem.7b00152, 2018.
- 25 Jang, J., Jang, M., Mui, W., Delcomyn, C. A., Henley, M. V., and Hearn, J. D.: Formation of active chlorine oxidants in saline-oxone aerosol, *Aerosol Sci Tech*, 44, 1018-1026, 2010.
- Jickells, T., An, Z., Andersen, K. K., Baker, A., Bergametti, G., Brooks, N., Cao, J., Boyd, P., Duce, R., and Hunter, K.: Global iron connections between desert dust, ocean biogeochemistry, and climate, *Science*, 308, 67-71, 2005.
- 30 Khorasani, A. E., Schroder, D. K., and Alford, T. L.: Optically Excited MOS-Capacitor for Recombination Lifetime Measurement, *Ieee Electr Device L*, 35, 986-988, 10.1109/Led.2014.2345058, 2014.
- 35 Kim, K. W., Kim, Y. J., and Oh, S. J.: Visibility impairment during Yellow Sand periods in the urban atmosphere of Kwangju, Korea, *Atmos Environ*, 35, 5157-5167, 2001.
- Krueger, B. J., Grassian, V. H., Laskin, A., and Cowin, J. P.: The transformation of solid atmospheric particles into liquid droplets through heterogeneous chemistry: Laboratory insights into the processing of calcium containing mineral dust aerosol in the troposphere, *Geophys Res Lett*, 30, Artn 1148. 10.1029/2002gl016563, 2003.
- 40 Li, J., Wang, Z. F., Zhuang, G., Luo, G., Sun, Y., and Wang, Q.: Mixing of Asian mineral dust with anthropogenic pollutants over East Asia: a model case study of a super-duststorm in March 2010, *Atmos Chem Phys*, 12, 7591-7607, 10.5194/acp-12-7591-2012, 2012.
- 45 Li, J., Jang, M., and Beardsley, R. L.: Dialkylsulfate formation in sulfuric acid-seeded secondary organic aerosol produced using an outdoor chamber under natural sunlight, *Environ Chem*, 2015.

- Li, W. J., and Shao, L. Y.: Chemical Modification of Dust Particles during Different Dust Storm Episodes, *Aerosol Air Qual Res*, 12, 1095-1104, 10.4209/aaqr.2011.11.0188, 2012.
- Liang, J. Y., and Jacobson, M. Z.: A study of sulfur dioxide oxidation pathways over a range of liquid water contents, pH values, and temperatures, *J Geophys Res-Atmos*, 104, 13749-13769, Doi 10.1029/1999jd900097, 1999.
- 5 Linsebigler, A. L., Lu, G. Q., and Yates, J. T.: Photocatalysis on TiO₂ Surfaces - Principles, Mechanisms, and Selected Results, *Chem Rev*, 95, 735-758, DOI 10.1021/cr00035a013, 1995.
- Liu, Y., Zhu, T., Zhao, D., and Zhang, Z.: Investigation of the hygroscopic properties of Ca(NO₃)₂ and internally mixed Ca(NO₃)₂/CaCO₃ particles by micro-Raman spectrometry, *Atmos Chem Phys*, 8, 7205-7215, 2008.
- 10 Mahowald, N. M., Baker, A. R., Bergametti, G., Brooks, N., Duce, R. A., Jickells, T. D., Kubilay, N., Prospero, J. M., and Tegen, I.: Atmospheric global dust cycle and iron inputs to the ocean, *Global biogeochemical cycles*, 19, 2005.
- 15 Michel, A. E., Usher, C. R., and Grassian, V. H.: Heterogeneous and catalytic uptake of ozone on mineral oxides and dusts: A Knudsen cell investigation, *Geophys Res Lett*, 29, 10-11-10-14, 10.1029/2002GL014896, 2002.
- Mochizuki, T., Kawamura, K., Aoki, K., and Sugimoto, N.: Long-range atmospheric transport of volatile monocarboxylic acids with Asian dust over a high mountain snow site, central Japan, *Atmos. Chem. Phys.*, 16, 14621-14633, 10.5194/acp-16-14621-2016, 2016.
- 20 Ndour, M., Nicolas, M., D'Anna, B., Ka, O., and George, C.: Photoreactivity of NO₂ on mineral dusts originating from different locations of the Sahara desert, *Phys Chem Chem Phys*, 11, 1312-1319, 10.1039/b806441e, 2009.
- Park, J., and Jang, M.: Heterogeneous photooxidation of sulfur dioxide in the presence of airborne mineral dust particles, *RSC Adv.*, 6, 58617-58627, 2016.
- 25 Park, J., Jang, M., and Yu, Z.: Heterogeneous Photo-oxidation of SO₂ in the Presence of Two Different Mineral Dust Particles: Gobi and Arizona Dust, *Environ Sci Technol*, 51, 9605-9613, 10.1021/acs.est.7b00588, 2017.
- Pradhan, M., Kyriakou, G., Archibald, A. T., Papageorgiou, A. C., Kalberer, M., and Lambert, R. M.: Heterogeneous uptake of gaseous hydrogen peroxide by Gobi and Saharan dust aerosols: a potential missing sink for H₂O₂ in the troposphere, *Atmos Chem Phys*, 10, 7127-7136, 10.5194/acp-10-7127-2010, 2010.
- 30 Sarwar, G., Fahey, K., Kwok, R., Gilliam, R. C., Roselle, S. J., Mathur, R., Xue, J., Yu, J., and Carter, W. P.: Potential impacts of two SO₂ oxidation pathways on regional sulfate concentrations: aqueous-phase oxidation by NO₂ and gas-phase oxidation by Stabilized Criegee Intermediates, *Atmos Environ*, 68, 186-197, 2013.
- 35 Sarwar, G., Simon, H., Fahey, K., Mathur, R., Goliff, W. S., and Stockwell, W. R.: Impact of sulfur dioxide oxidation by Stabilized Criegee Intermediate on sulfate, *Atmos Environ*, 85, 204-214, 2014.
- 40 Scheuvens, D., and Kandler, K.: On Composition, Morphology, and Size Distribution of Airborne Mineral Dust, in: *Mineral Dust: A Key Player in the Earth System*, edited by: Knippertz, P., and Stuut, J.-B. W., Springer Netherlands, Dordrecht, 15-49, 2014.
- Schulz, M., Prospero, J. M., Baker, A. R., Dentener, F., Ickes, L., Liss, P. S., Mahowald, N. M., Nickovic, S., Garcia-Pando, C. P., Rodriguez, S., Sarin, M., Tegen, I., and Duce, R. A.:

- Atmospheric Transport and Deposition of Mineral Dust to the Ocean: Implications for Research Needs, *Environ Sci Technol*, 46, 10390-10404, 10.1021/es300073u, 2012.
- 5 Sokolik, I., Winker, D., Bergametti, G., Gillette, D., Carmichael, G., Kaufman, Y., Gomes, L., Schuetz, L., and Penner, J.: Introduction to special section: Outstanding problems in quantifying the radiative impacts of mineral dust, *Journal of Geophysical Research: Atmospheres*, 106, 18015-18027, 2001.
- Sokolik, I. N., and Toon, O. B.: Direct radiative forcing by anthropogenic airborne mineral aerosols, *Nature*, 381, 681-683, 1996.
- 10 Tang, M. J., Whitehead, J., Davidson, N. M., Pope, F. D., Alfarra, M. R., McFiggans, G., and Kalberer, M.: Cloud condensation nucleation activities of calcium carbonate and its atmospheric ageing products, *Phys. Chem. Chem. Phys.*, 17, 32194-32203, 10.1039/c5cp03795f, 2015.
- Tang, M. J., Cziczo, D. J., and Grassian, V. H.: Interactions of Water with Mineral Dust Aerosol: Water Adsorption, Hygroscopicity, Cloud Condensation, and Ice Nucleation, *Chem Rev*, 15 116, 4205-4259, 10.1021/acs.chemrev.5b00529, 2016.
- Tang, M. J., Huang, X., Lu, K. D., Ge, M. F., Li, Y. J., Cheng, P., Zhu, T., Ding, A. J., Zhang, Y. H., Gligorovski, S., Song, W., Ding, X., Bi, X. H., and Wang, X. M.: Heterogeneous reactions of mineral dust aerosol: implications for tropospheric oxidation capacity, *Atmos Chem Phys*, 17, 11727-11777, 10.5194/acp-17-11727-2017, 2017.
- 20 Tegen, I., and Schepanski, K.: The global distribution of mineral dust, *IOP Conference Series: Earth and Environmental Science*, 2009, 012001,
- Textor, C., Schulz, M., Guibert, S., Kinne, S., Balkanski, Y., Bauer, S., Berntsen, T., Berglen, T., Boucher, O., Chin, M., Dentener, F., Diehl, T., Easter, R., Feichter, H., Fillmore, D., Ghan, S., Ginoux, P., Gong, S., Grini, A., Hendricks, J., Horowitz, L., Huang, P., Isaksen, I., 25 Iversen, I., Kloster, S., Koch, D., Kirkevåg, A., Kristjansson, J. E., Krol, M., Lauer, A., Lamarque, J. F., Liu, X., Montanaro, V., Myhre, G., Penner, J., Pitari, G., Reddy, S., Seland, Ø., Stier, P., Takemura, T., and Tie, X.: Analysis and quantification of the diversities of aerosol life cycles within AeroCom, *Atmos. Chem. Phys.*, 6, 1777-1813, 10.5194/acp-6-1777-2006, 2006.
- 30 Thompson, T. L., and Yates, J. T.: Surface science studies of the photoactivation of TiO₂ new photochemical processes, *Chem Rev*, 106, 4428-4453, 2006.
- Underwood, G. M., Song, C. H., Phadnis, M., Carmichael, G. R., and Grassian, V. H.: Heterogeneous reactions of NO₂ and HNO₃ on oxides and mineral dust: A combined laboratory and modeling study, *J Geophys Res-Atmos*, 106, 18055-18066, Doi 35 10.1029/2000jd900552, 2001.
- Uno, I., Osada, K., Yumimoto, K., Wang, Z., Itahashi, S., Pan, X. L., Hara, Y., Yamamoto, S., and Nishizawa, T.: Importance of Long-Range Nitrate Transport Based on Long-Term Observation and Modeling of Dust and Pollutants over East Asia, *Aerosol Air Qual Res*, 17, 3052-3064, 10.4209/aaqr.2016.11.0494, 2017.
- 40 Usher, C. R., Michel, A. E., and Grassian, V. H.: Reactions on mineral dust, *Chem Rev*, 103, 4883-4939, 10.1021/cr020657y, 2003a.
- Usher, C. R., Michel, A. E., Stec, D., and Grassian, V. H.: Laboratory studies of ozone uptake on processed mineral dust, *Atmos Environ*, 37, 5337-5347, 10.1016/j.atmosenv.2003.09.014, 2003b.

- Valle-Diaz, C. J., Torres-Delgado, E., Colon-Santos, S. M., Lee, T., Collett, J. L., McDowell, W. H., and Mayol-Bracero, O. L.: Impact of Long-Range Transported African Dust on Cloud Water Chemistry at a Tropical Montane Cloud Forest in Northeastern Puerto Rico, *Aerosol Air Qual Res*, 16, 653-664, 10.4209/aaqr.2015.05.0320, 2016.
- 5 Vlasenko, A., Sjogren, S., Weingartner, E., Stemmler, K., Gäggeler, H. W., and Ammann, M.: Effect of humidity on nitric acid uptake to mineral dust aerosol particles, *Atmos. Chem. Phys.*, 6, 2147-2160, 10.5194/acp-6-2147-2006, 2006.
- 10 Wang, Z., Pan, X. L., Uno, I., Li, J., Wang, Z. F., Chen, X. S., Fu, P. Q., Yang, T., Kobayashi, H., Shimizu, A., Sugimoto, N., and Yamamoto, S.: Significant impacts of heterogeneous reactions on the chemical composition and mixing state of dust particles: A case study during dust events over northern China, *Atmos Environ*, 159, 83-91, 10.1016/j.atmosenv.2017.03.044, 2017.
- 15 Wexler, A. S., and Clegg, S. L.: Atmospheric aerosol models for systems including the ions H^+ , NH_4^+ , Na^+ , SO_4^{2-} , NO_3^- , Cl^- , Br^- , and H_2O , *Journal of Geophysical Research: Atmospheres*, 107, 4207, 2002.
- 20 Yang, W. W., Zhang, J. H., Ma, Q. X., Zhao, Y., Liu, Y. C., and He, H.: Heterogeneous Reaction of SO_2 on Manganese Oxides: the Effect of Crystal Structure and Relative Humidity, *Sci Rep-Uk*, 7, ARTN 4550. 10.1038/s41598-017-04551-6, 2017.
- 25 Yu, Z., Jang, M., and Park, J.: Modeling atmospheric mineral aerosol chemistry to predict heterogeneous photooxidation of SO_2 , *Atmos Chem Phys*, 17, 10001-10017, 2017.
- Zender, C. S., Miller, R. L. R. L., and Tegen, I.: Quantifying mineral dust mass budgets: Terminology, constraints, and current estimates, *Eos, Transactions American Geophysical Union*, 85, 509-512, doi:10.1029/2004EO480002, 2004.
- Zhong, M., and Jang, M.: Dynamic light absorption of biomass-burning organic carbon photochemically aged under natural sunlight, *Atmos Chem Phys*, 14, 1517-1525, 2014.

Table 1. Outdoor chamber experimental conditions for NO_x and SO₂ oxidation in the presence of Gobi Desert dust (GDD) particles and Arizona Test dust (ATD) particles.

Exp. No	Purpose	Type of particles	Mass conc. of particles ^a (µg m ⁻³) ^b	RH ^b (%)	Temp ^b . (°C)	Initial NO/NO ₂ conc. (ppb) ^b	Initial SO ₂ conc. (ppb) ^b	Initial O ₃ conc. (ppb) ^b
10/6/2017	High and low NO _x with SO ₂	GDD	337.3	13.9-91.8	293.9-319.3	22.1/123.1	93.9	4.1
		GDD	375.3	21.9-95.6	294.3-320.3	6.1/37.1	98.2	6.0
17/9/2017	GDD vs. ATD with NO _x	ATD	334.0	14.2-50.9	293.6-319.4	19.1/108.1	N.A. ^c	3.6
		GDD	408.1	21.0-61.6	294.0-318.9	17.1/99.1	N.A. ^c	2.8

^a The mass concentrations of GDD and ATD particles were calculated from the SMPS data combined with OPC data. The density of dust particles is 2.65 g cm⁻³ and the particle size distribution was calculated up to 3 µm.

^b The errors associated with NO, NO₂, and O₃ were ± 12.5%, ± 6.9%, and ± 0.2%, respectively. The error associated with dust mass were ± 6% based on SMPS and OPC data. The accuracy of the measurement of RH and temperature were ± 5 % and ± 0.5 K, respectively.

^c N.A.: not applicable (no SO₂ injection).

10

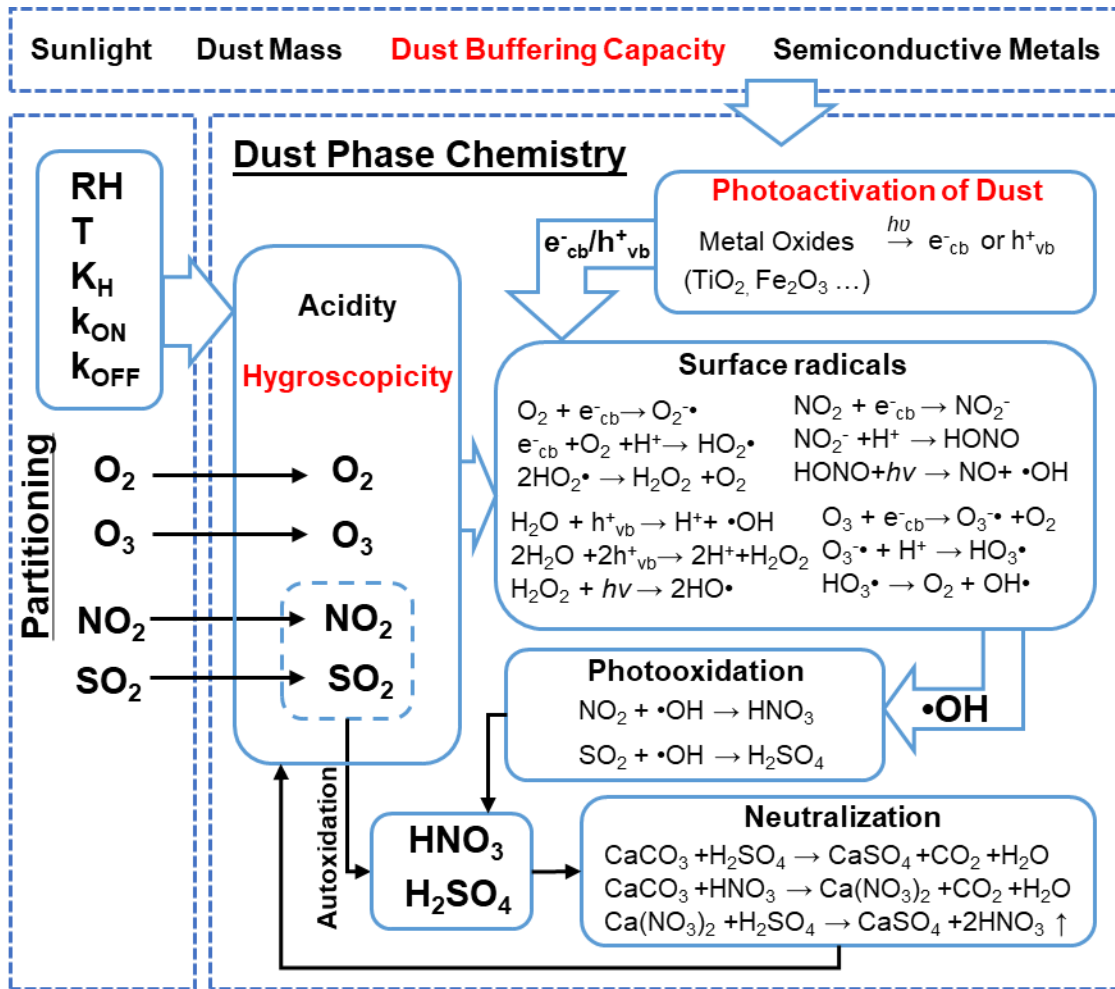


Figure 1. The overall schematic of dust phase chemistry in the AMAR model.

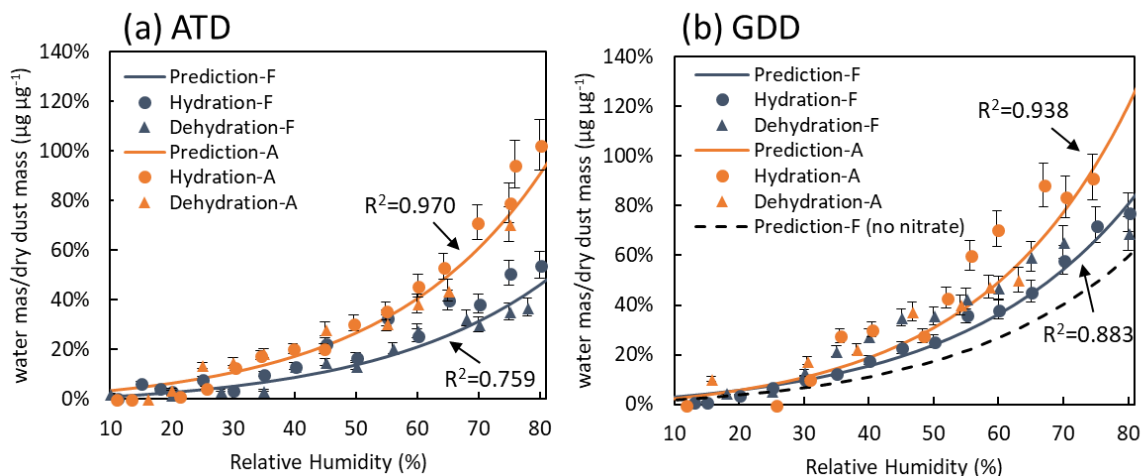


Figure 2. The fraction of water mass relative to the dry dust mass for fresh and photochemically aged (a) ATD and (b) GDD particles as a function of the relative humidity from 10% to 85%. “F” and “A” represent the fresh and aged dust particles, respectively. The water content for fresh GDD with theoretically no indigenous nitrate is also predicted and shown. The aged dust samples were collected from dust particles that were photochemically aged in the presence of NO_x . The estimated nitrate concentrations for fresh and aged ATD are $0.001 \mu\text{g } \mu\text{g}^{-1}$ and $0.011 \mu\text{g } \mu\text{g}^{-1}$, respectively. The estimated nitrate concentrations for fresh and aged GDD are $0.007 \mu\text{g } \mu\text{g}^{-1}$ and $0.02 \mu\text{g } \mu\text{g}^{-1}$, respectively. The error bars were estimated from the uncertainties in the FTIR absorbance measurements of the O-H band.

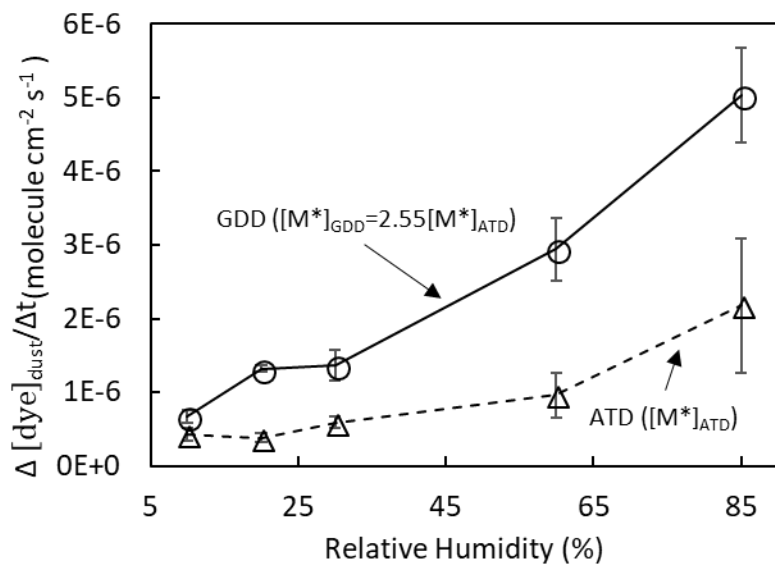


Figure 3. The dye degradation rate in the presence of ATD or GDD particles measured using a UV flow chamber under RH levels ranging from 10% to 85%. As a control, the photodegradation of malachite green in the absence of dust was measured, but the degradation was negligible. The error bars represent the standard deviations.

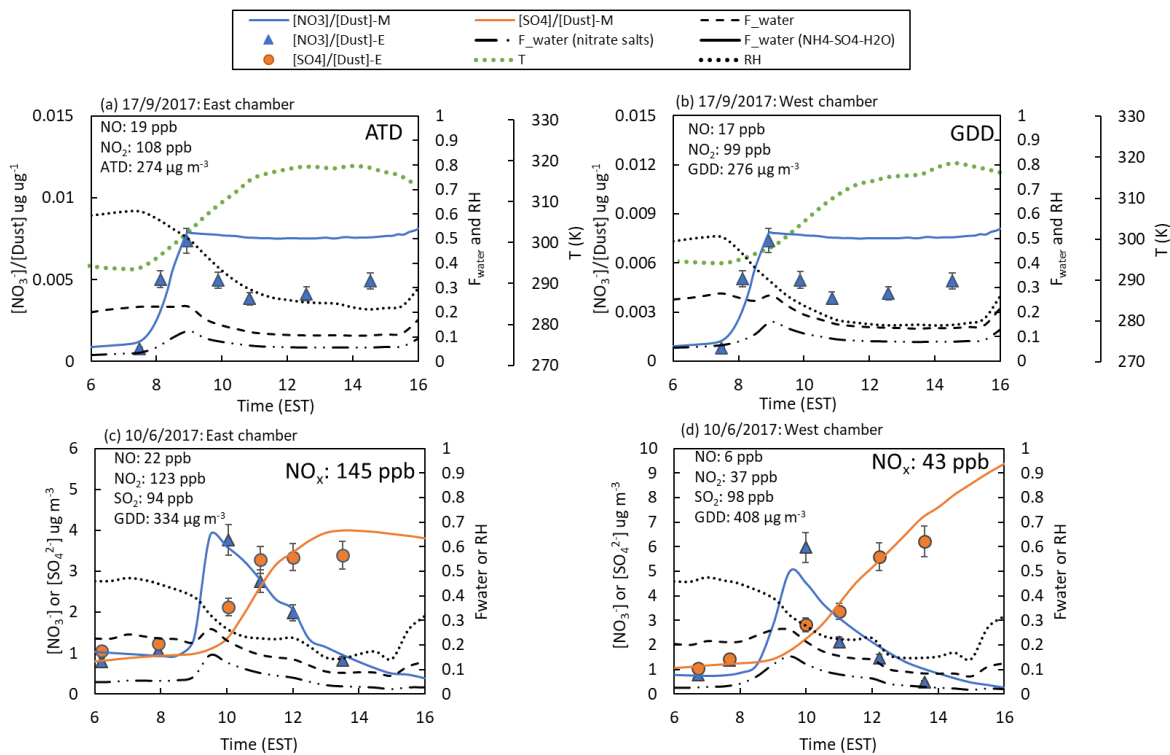


Figure 4. Simulation of the outdoor chamber data using the AMAR model for (A) ATD (2017-09-17) and (B) GDD (2017-09-17) particles in the presence of NO_x and SO_2 oxidation on GDD particles in the presence of (C) high NO_x (2017-06-10) and (D) low NO_x (2017-06-10) concentrations. $F_{\text{water}}(\text{nitrate salts})$ and $F_{\text{water}}(\text{NH}_4^+-\text{SO}_4^{2-}-\text{H}_2\text{O})$ are the second and third terms in Eq. (1) and represent the additional absorbed water by alkaline nitrate salts and the ammonium sulfate system, respectively. The simulation result was not correct for particle loss.

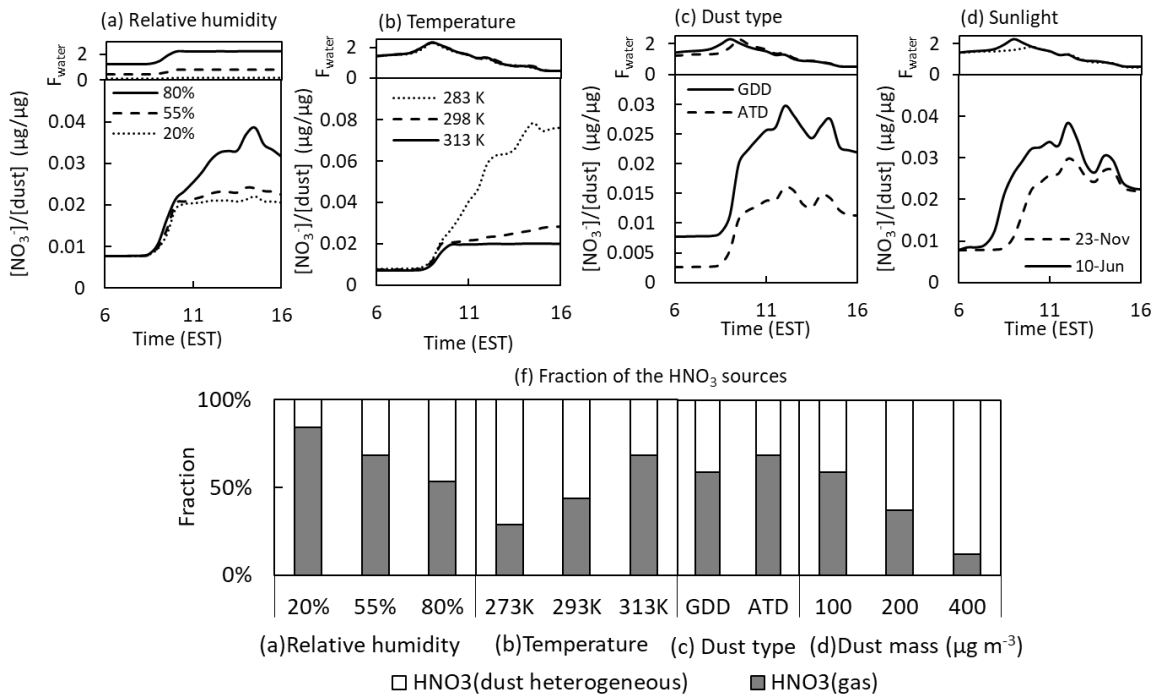


Figure 5. Sensitivity of the model nitrate prediction to the (a) relative humidity at 20%, 55% and 80%; (b) temperature at 273 K, 293 K and 313 K; (c) ATD vs. GDD particles; (d) concentration of GDD at 100, 200 and 400 $\mu\text{g m}^{-3}$; and (e) sunlight profile on 23 November 2017 vs. 10 June 2017. The fraction of the HNO₃ sources formed from the gas-phase reaction and dust-phase heterogeneous reaction to the total HNO₃ is shown in (f). The simulation was conducted with 100 $\mu\text{g m}^{-3}$ of initial GDD particles, 40 ppb of initial NO_x (NO:NO₂=1:1), 2 ppb of initial O₃ and 10 ppb isoprene under ambient environmental conditions on 23 November 2017. The simulation was performed without considering particle loss.

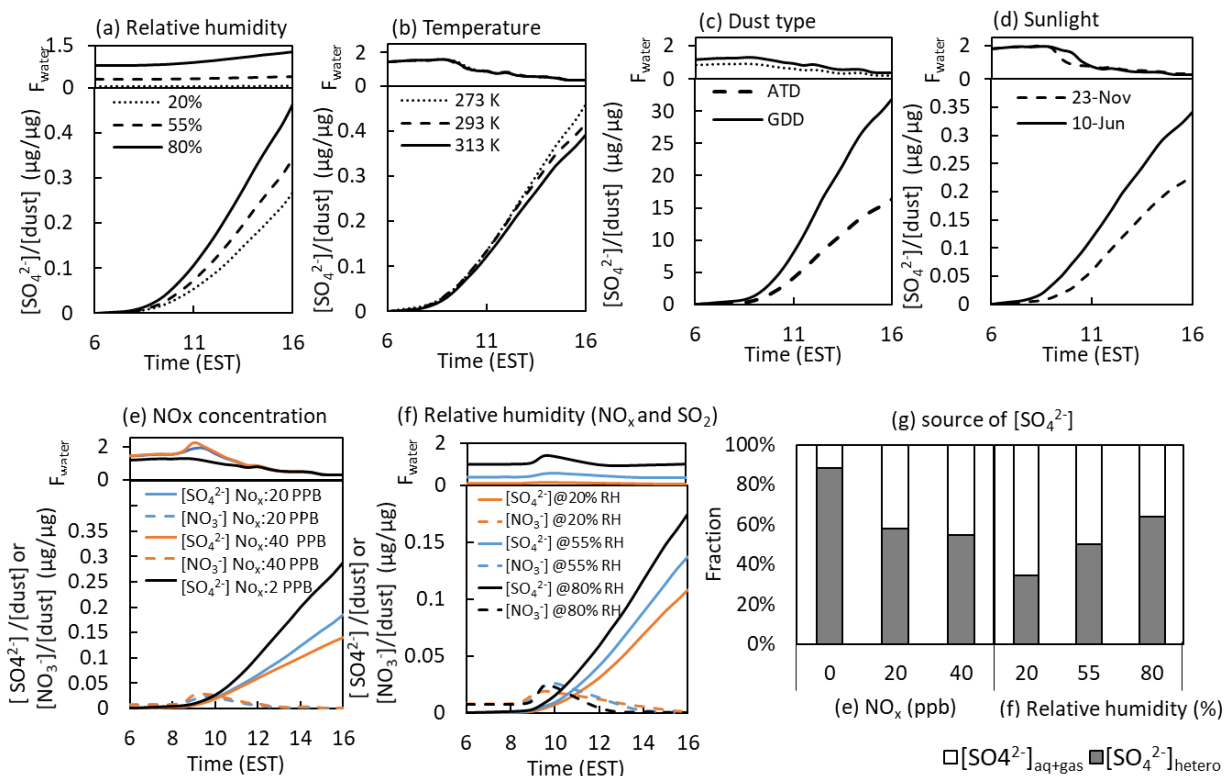
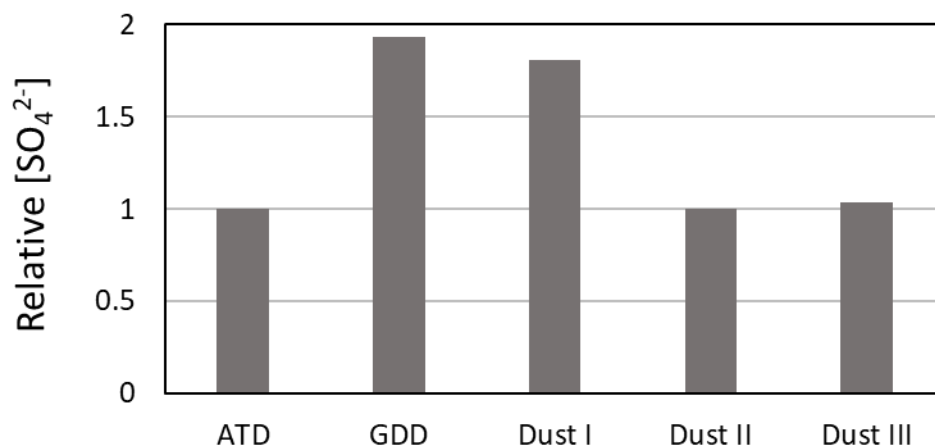


Figure 6. Sensitivity of the model predicted sulfate concentration to the (a) relative humidity at 20%, 55% and 80%; (b) temperature at 273 K, 293 K and 313 K; (c) sunlight profile (23 November 2017 vs. 10 June 2017); (d) dust type (ATD vs. GDD); (e) initial concentration of NO_x (0, 20 and 40 ppb); and (f) relative humidity (20%, 55% and 80%) in the presence of 20 ppb of NO_x. The fraction of sulfate from the gas phase and non-dust aqueous phase ($[SO_4^{2-}]_{aq+gas}$) and the heterogeneously formed sulfate in the dust phase ($[SO_4^{2-}]_{hetero}$) relative to the total sulfate is shown in (g). The dust-phase nitrate and water content were also predicted. For the sensitivity test, the simulation was conducted with $100 \mu\text{g m}^{-3}$ of initial GDD particles, 40 ppb of initial SO₂, 2 ppb of initial O₃ and 10 ppb isoprene under ambient environmental conditions on 23 November 2017. The simulation was performed without considering particle loss.



Dust characteristics	Dust type				
	ATD	GDD	Dust I	Dust II	Dust III
Photoactivation ($[M^*]_{dust}$, Eq. 2)	1	2.55	2.55	1	1
Buffering capacity (Sect. 3.4)	1	2	1	1	2
First term of F_{water} (Eq. 1)	1	1	1	1.5	1

Figure 7. The analysis of the influential parameters associated with dust characteristics to form sulfate. The relative concentration of sulfate is predicted using AMAR in the presence of different types of dust including ATD, GDD and three types of artificial dust (Dust I, II and III). The variation of dust type is determined *via* three major aspects: photoactivation capability of dust linked to $[M^*]$ in Eq. 2 (Sect. 3.3), the buffering capacity of dust (Sect. 3.4) and F_{water} in Eq. 1 (Sect. 3.2). Dust I, II and III are artificially formulated to analyse how the three dust properties can influence the sulfate formation. ATD is used as a reference dust. The three parameters of GDD, which were obtained from experimental data, are scaled to those of ATD. For analysis, the simulation is conducted with $100 \mu\text{g m}^{-3}$ of initial dust particles, 40 ppb of initial SO_2 , 2 ppb of initial O_3 and 10 ppb isoprene under ambient environmental condition on 23 November 2017. The simulation was performed without considering the particle loss to the chamber wall.