



1 **Two years online measurement of fine particulate nitrate in western**  
2 **Yangtze River Delta: Influences of thermodynamics and N<sub>2</sub>O<sub>5</sub> hydrolysis**

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11 **Abstract.**

12 Particulate nitrate contributes a large fraction of secondary aerosols. Despite  
13 understanding of its important role in regional air quality and global climate, long-term  
14 continuous measurements are rather limited in China. In this study, we conducted online  
15 measurement of PM<sub>2.5</sub> nitrate for two years from March 2014 to February 2016 using  
16 the Monitor for Aerosols and Gases in ambient Air (MARGA) in the western Yangtze  
17 River Delta (YRD), eastern China, and investigate the main factors that influenced its  
18 temporal variations and formation pathways. Compared to other sites in China, an  
19 overall high concentration of particulate nitrate was observed with a mean value of 15.8  
20  $\mu\text{g m}^{-3}$  (0.5 to 92.6  $\mu\text{g m}^{-3}$ ). Nitrate on average accounted for 32% of the total mass of  
21 water-soluble ions and the proportion increased with PM loading, indicating that nitrate  
22 is a major driver of haze pollution episodes in this region. Sufficient ammonia drove  
23 most nitrate into the particle phase in the form of ammonium nitrate. A typical seasonal  
24 cycle of nitrate was observed with the concentrations in winter on average two times  
25 higher than those in summer mainly due to different meteorological conditions. In  
26 summer, the diurnal variation of particulate nitrate was determined by the  
27 thermodynamic equilibrium, resulting in a much lower concentration during daytime  
28 despite of a considerable photochemical production. Air masses from polluted YRD



29 and biomass burning region contributed to the high nitrate concentration during summer.  
30 In winter, particulate nitrate didn't reveal an evident diurnal variation. Regional  
31 transport from northern China played an important role in enhancing nitrate  
32 concentration. Eighteen nitrate episodes were selected to understand the processes that  
33 drive the formation of high concentration of nitrate. Rapid nitrate formation was  
34 observed during the pre-episode (the day before nitrate episode day) nights, and  
35 dominated the increase of total water-soluble ions. Calculated nitrate from  $\text{N}_2\text{O}_5$   
36 hydrolysis was highly correlated to and accounted for 80 percent of the observed nitrate,  
37 suggesting that  $\text{N}_2\text{O}_5$  hydrolysis was a major contributor to the nitrate episodes. Our  
38 results suggested that rapid formation of nitrate could be a main cause for extreme  
39 aerosol pollution events in YRD during winter, and illustrated the urgent needs to  
40 control the  $\text{NO}_x$  emission.

#### 41 **1. Introduction**

42 Particulate nitrate ( $\text{NO}_3^-$ ), as a major aerosol component in the atmosphere, reduces  
43 atmospheric visibility (Charlson and Heintzenberg, 1995), influences human health,  
44 alters radiative forcing and hence influences regional even global climate (IPCC, 2013).  
45 Compared to the sulfate, nitrate has a larger scattering albedo under low RH conditions  
46 that cause a stronger influence on visibility (Lei and Wuebbles, 2013). High  
47 concentration of particulate nitrate had been demonstrated to be one of the major  
48 reasons for the frequent occurrence of haze episodes in China (Wang and Zhang, 2009;  
49 Wen et al., 2015; Wang et al., 2017). In recent decades, Chinese government started to  
50 control emissions of air pollutants with special effort on the  $\text{SO}_2$  reduction. This resulted  
51 in a remarkable decrease of ambient  $\text{SO}_2$  and sulfate concentrations after 2006 (van der  
52 A et al., 2017; Wang et al., 2017). However, particulate nitrate, as well as its proportion  
53 in PM, showed increasing trends due to the strong emission of nitrogen oxides ( $\text{NO}_x$ )  
54 (Lei and Wuebbles, 2013; Yang et al., 2017).

55 Particulate nitrate can be formed from multiple pathways. Gas phase reaction of  
56  $\text{NO}_2$  and OH radical is one major pathway to form nitric acid ( $\text{HNO}_3$ ) (Calvert and



57 Stockwell, 1983), which subsequently reacts with ammonia ( $\text{NH}_3$ ) to produce  
58 ammonium nitrate ( $\text{NH}_4\text{NO}_3$ ). As typical photochemical processes, these reactions  
59 dominate daytime nitrate formation, and have been widely investigated in both field  
60 and modelling studies (Sharma et al., 2007; Petetin et al., 2016). Heterogeneous uptake  
61 of the photochemical formed nitric acid by alkali compounds, e.g. dust and sea salt  
62 particles, is also a considerable pathway to form nitrate in some regions (Bian et al.,  
63 2014). During nighttime, the hydrolysis of dinitrogen pentoxide ( $\text{N}_2\text{O}_5$ ) is believed to  
64 be the dominate pathway to form particulate nitrate.  $\text{N}_2\text{O}_5$  is an important reactive  
65 nitrogen species in the polluted troposphere (Brown and Dube., 2007; Osthoff et al.,  
66 2006; Li et al., 2017; Brown et al., 2003; Brown and Stutz, 2012) and accumulates via  
67 the reversible reaction between  $\text{NO}_2$  and  $\text{NO}_3$  radical produced from the reaction of  
68  $\text{NO}_2$  with  $\text{O}_3$ . Due to the rapid photolysis of  $\text{NO}_3$  radical,  $\text{N}_2\text{O}_5$  concentration during  
69 daytime is rather low that its contribution to nitrate can be ignored. While during  
70 nighttime,  $\text{N}_2\text{O}_5$  concentration can be up to ppb level, and form nitric acid by reaction  
71 with water vapor, or particulate nitrate directly by heterogeneous hydrolysis on the wet  
72 surface (Wang et al., 2017; Wen et al., 2018; Thornton et al., 2003). In China, the  
73 pollution episodes with high nitrate concentrations mostly occurred in winter, during  
74 which the photochemical production of nitrate should be overall weak.  $\text{N}_2\text{O}_5$  hydrolysis  
75 thus has the potential to be the crucial contributor, however there is still a lack of  
76 observational evidences.

77 Collecting particulate matter on a filter with subsequent ion chromatography  
78 analysis in laboratories is the conventional method to measure the concentration of  
79 particulate nitrate. The un-denude filter pack system, which is most-widely used, can  
80 suffer from both positive artifacts by absorbing gas-phase nitric acid, and negative  
81 artifacts by the evaporation of ammonium nitrate (Nie et al., 2010; Pathak and Wu,  
82 2009; Wang et al., 2010). A denuder system can minimize these sampling artifacts by  
83 adding denuders to remove the interfering gases and back-up filters to collect the  
84 evaporated vapors (John et al., 1988). However, the operation of such a denuder system



85 is super labor intensive and thus not widely used. In addition, the poor time-resolution  
86 of filter-based measurement can limit our understanding on the formation and chemical  
87 evolution of the particulate nitrate. To overcome these shortcomings, several  
88 continuous and semi-continuous techniques have been developed based on an online  
89 denude-IC system (e.g. the ambient ion monitor (AIM), the gas and aerosol collector  
90 ion chromatography (GAC-IC), the particle-into-liquid sample ion chromatography  
91 (PILS-IC) and MARGA), as well as mass spectrometry (e.g. AMS). Pathak et al. (2011)  
92 applied an AIM instrument in Beijing and Shanghai for one month and found that the  
93 heterogeneous hydrolysis of  $N_2O_5$  contributed 50%~100% of the nighttime nitrate  
94 formation. Xue et al. (2013) deployed a PILS-IC system in Hong Kong for less than a  
95 month and showed a more active nitrate formation during PM episode than normal days.  
96 Wen et al. (2015) used a MARGA instrument in Yucheng, North China during summer  
97 and emphasized the important roles of  $O_3$  and  $NH_3$  on nitrate formation. Yang et al.  
98 (2017) carried out field observation with ACSM in Beijing for half a month and pointed  
99 out the importance of aerosol nitrate in haze formation. However, despite of an  
100 increasing number of studies using online techniques, continuous measurements with  
101 more than one-year period are still very limited.

102 The Yangtze River Delta (YRD) located in the eastern China, with megacities  
103 including Shanghai, Nanjing and Suzhou etc., has suffered from heavy particulate  
104 matter pollution and photochemical pollution (Ding et al., 2013bc; Wang et al., 2016bc;  
105 Wang et al., 2016a). Previous studies indicated an important role of nitrate in the  
106 pollution episodes (Hua et al., 2015; Du et al., 2011; Yang et al., 2017). For example,  
107 Zhang et al. (2015) carried out an observation with ACSM in urban Nanjing during  
108 summer and autumn, and found nitrate and organic aerosols dominated the  $PM_{10}$   
109 composition. Shi et al. (2014) used MARGA instrument in Shanghai for months and  
110 found an increasing contribution of nitrate to  $PM_{10}$  mass during pollution periods.  
111 Wang et al. (2016b) reported temporal variation and transport of  $PM_{2.5}$  water soluble  
112 ions, including nitrate, in an urban site in Shanghai based on three-year continuous



113 measurement using MARGA. However, detailed investigation on the possible  
114 mechanisms governing nitrate behaviors during haze pollution is still rare.

115 In this study, we present a 2-year continuous measurement of particulate nitrate  
116 using MARGA at a rural site in Nanjing, a megacity in the western YRD region, with  
117 the target to get a comprehensive understanding of particulate nitrate behaviors and  
118 investigate the processes affecting nitrate in haze episodes. We first conducted general  
119 statistical analysis of particulate nitrate and characterized seasonal variation and diurnal  
120 pattern. A thermodynamic model was then applied to investigate the gas-particle  
121 partition of nitrate. The influence of air masses was also investigated by conducting  
122 backward Lagrangian dispersion modelling. Finally, we selected eighteen nitrate  
123 episodes and investigated the main processes influencing their evolution.

## 124 2. Methodology

### 125 2.1. Sample site and instrumentation

126 The SORPES station (118°57'E, 32°07'N) was located on the top of a small hill (40 m  
127 above sea level) in the Xianlin campus of Nanjing University located in the outskirts of  
128 Nanjing, China. The station is an ideal receptor of air masses from the YRD with little  
129 influence of local emissions and urban pollution from Nanjing. Detailed description can  
130 be found in previous studies (Ding et al., 2013c; Ding et al., 2016a).

131 The measurement was conducted from March 2014 to Feb 2016. Hourly  
132 concentrations of water soluble gases of HCl, HNO<sub>3</sub>, HONO, SO<sub>2</sub> and NH<sub>3</sub>, and water-  
133 soluble ions in PM<sub>2.5</sub>, including Cl<sup>-</sup>, NO<sub>3</sub><sup>-</sup>, SO<sub>4</sub><sup>2-</sup>, NH<sub>4</sub><sup>+</sup>, Na<sup>+</sup>, K<sup>+</sup>, Ca<sup>2+</sup>, and Mg<sup>2+</sup>, were  
134 measured with a Monitor for Aerosols and Gases in ambient Air (MARGA, designed  
135 and manufactured by Applikon Analytical B.V., the Netherlands) employed in  
136 connection with a Thermo PM<sub>2.5</sub> cyclone inlet. The sampling system was comprised of  
137 two parts: A Wet Rotating Denuder for gases and a Steam Jet Aerosol Collector for  
138 aerosols, working at an air flow of 1 m<sup>3</sup> h<sup>-1</sup> (ten Brink et al., 2007; Rumsey et al., 2014).  
139 After each hour's collection, the samples were analyzed using ion chromatography. The  
140 instrument was calibrated on an hourly basis using internal standard liquid (bromide



141 lithium), ensuring a stable and reliable ion chromatograph. Concentrations of all aerosol  
142 ions and gases have a precision of  $0.001 \mu\text{g m}^{-3}$  (Xie et al., 2015). The  $\text{PM}_{2.5}$  ion dataset  
143 from the MARGA provided more than 15000 hourly samples over the 24 months of  
144 measurements considering points where  $\text{NO}_3^-$ ,  $\text{NH}_4^+$  and  $\text{SO}_4^{2-}$  were all available. Trace  
145 gases (i.e.,  $\text{O}_3$ ,  $\text{SO}_2$ ,  $\text{NO}_x$ ,  $\text{NO}$ ) and  $\text{PM}_{2.5}$  mass concentrations were also measured at  
146 SORPES (Ding et al., 2013c; Nie et al., 2015; Ding et al., 2016a), together with  
147 meteorological data including wind speed/direction, temperature, and relative humidity.

## 148 2.2. Thermodynamic constants and ISORROPIA II

149 Formation of ammonium nitrate involves an equilibrium reaction between the gas phase  
150  $\text{NH}_3$  and  $\text{HNO}_3$ , and particle phase  $\text{NH}_4\text{NO}_3$ . The gas-to-particle partitioning is  
151 temperature dependent, and the equilibrium constant can be calculated as follows (units  
152  $\text{mol}^2\text{kg}^{-2}\text{atm}^{-2}$ ) (Sun et al., 2011; Seinfeld and Pandis, 2006):

$$153 \quad K = k_{298} \exp(a(298/T-1) + b[1 + \ln(298/T) - 298/T]) \quad (1)$$

$$154 \quad K_{298} = 3.5 \times 10^{16} (\text{atm}^{-2}), \quad a = 75.11, \quad b = -13.5. \quad (2)$$

155 The dissociation constant,  $K_p(T)$ , is the value to examine the product of the partial  
156 pressures of  $\text{NH}_3$  and  $\text{HNO}_3$  in ideal equilibrium state at a specific temperature and can  
157 be calculated as follows (units:  $\text{ppb}^2$ ) (Sharma et al., 2007; Seinfeld and Pandis, 2006):

$$158 \quad \ln K_p = 84.6 - 24220/T - 6.4 \ln(T/298) \quad (T \text{ is temperature in Kelvin}) \quad (3)$$

159 Note that the constant is quite sensitive to temperature changes. Lower values of  $K_p$   
160 correspond to lower equilibrium values of the  $\text{NH}_3$  and  $\text{HNO}_3$  gas phase concentrations,  
161 shifting the equilibrium of the system toward the aerosol phase.

162 ISORROPIA II (available at <http://isorroopia.eas.gatech.edu/>) is a thermodynamic  
163 model used commonly in inorganic aerosol research (Fountoukis and Nenes, 2007). To  
164 analyze gas-into-particulate pathway for nitrate formation,  $\text{HNO}_3$  was modeled with  
165 ISORROPIA II run in forward model iteratively (Pusede et al., 2016; Fountoukis and  
166 Nenes, 2007). ISORROPIA II was initialized as  $[\text{NO}_3^- + \text{HNO}_3]_{\text{total}} = [\text{NO}_3^-]_{\text{aerosol}}$ .  
167 Calculated  $\text{HNO}_3(\text{g})$  was added back to  $[\text{NO}_3^- + \text{HNO}_3]$ , while we always use  
168  $[\text{NH}_4^+_{\text{aerosol}} + \text{NH}_3(\text{g})]$  as input total ammonium. ISORROPIA II was solved iteratively



169 until output  $\text{NO}_3^-$  changed by  $< 2\%$  by mass. The phase state was set as metastable. We  
170 assume that gases and aerosol are in equilibrium, that aerosols are homogeneous and  
171 internally mixed, and that unaccounted-for factors do not influence the thermodynamics  
172 of system (Vayenas et al., 2005).

### 173 2.3. Lagrangian Dispersion Modeling

174 To help understand the influence of air masses, backward Lagrangian particulate  
175 dispersion modeling (LPDM) was carried out based on a method developed and  
176 evaluated by Ding (Ding et al., 2013a). The LPDM was conducted using the Hybrid  
177 Single-Particulate Lagrangian Integrated Trajectory model developed in the Air  
178 Resource Laboratory (ARL) of the National Oceanic and Atmospheric Administration  
179 using the ARL format Global Data Assimilation System data. The model calculates the  
180 position of particulates by mean wind and a turbulence transport component after they  
181 are released at the source point for a backward simulation. For each hour, 3000  
182 particulates were released at 100 m altitude over the site and were traced backward for  
183 a 3-day period. The hourly position of each particulate was calculated using a 3-D  
184 particulate, i.e., horizontal and vertical, method. The residence time at 100 m altitude,  
185 i.e., foot-print “retroplume”, which represents the distribution of the surface probability  
186 or residence time of the simulated air mass, was used to understand the contribution  
187 from potential source regions (Ding et al., 2013ac; Shen et al., 2018).

### 188 2.4. Steady-state predictions

189 Based on their short lifetimes, the concentrations of the  $\text{NO}_3$  radical and  $\text{N}_2\text{O}_5$  can be  
190 predicted by steady-state calculations due to lack of measurement data (Osthoff et al.,  
191 2006). The formation and loss of  $\text{N}_2\text{O}_5$  associated with a series of chemical reactions  
192 are listed in Table 1. For the heterogeneous processes, we used 0.004 and 0.03 as the  
193 uptake coefficients of the  $\text{NO}_3$  radical and  $\text{N}_2\text{O}_5$  ( $\gamma\text{NO}_3$  and  $\gamma\text{N}_2\text{O}_5$ ), respectively  
194 (Aldener et al., 2006; Wen et al., 2015; Knopf et al., 2011; Brown et al., 2006). Due to  
195 the lack of VOC measurement data, the total reaction rates of  $\text{NO}_3$  with VOCs were  
196 assumed to be equal to the  $\text{NO}_3$  loss rate caused by the heterogeneous hydrolysis of



197  $\text{N}_2\text{O}_5$  (Aldener et al., 2006; Wen et al., 2015). The corresponding rate constant can be  
198 found in Master Chemical Mechanism (MCM version 3.1,  
199 <http://mcm.leeds.ac.uk/MCM/>).

### 200 3. Results and discussion

#### 201 3.1. Overall results

202 A MARGA was deployed to continuously measure the eight water-soluble ions (WSI)  
203 of  $\text{PM}_{2.5}$  and several gas-phase species from March 2014 to February 2016 with the  
204 time resolution of 1-hour. Nitrate ( $\text{NO}_3^-$ ), sulfate ( $\text{SO}_4^{2-}$ ) and ammonium ( $\text{NH}_4^+$ ) were  
205 the major components with the two-year averaged concentrations of 15.8 ( $\pm 13.4$ ), 15.3  
206 ( $\pm 10.6$ ) and 10.4 ( $\pm 7.6$ )  $\mu\text{g m}^{-3}$ , respectively. In the present study, we focused on nitrate,  
207 and discussed the temporal variation and its association with physicochemical  
208 processes. The concentration of  $\text{PM}_{2.5}$  nitrate changed largely from 0.5 to 92.6  $\mu\text{g m}^{-3}$   
209 during the measurement period, and accounted for 3% to 58% of total WSI ( $=45.7 +$   
210  $30\mu\text{g m}^{-3}$ ). The highest hourly nitrate concentration occurred on December 23, 2015  
211 together with high concentrations of sulfate (65.5  $\mu\text{g m}^{-3}$ ), and ammonium (56.8  $\mu\text{g m}^{-3}$ ).  
212 Heavy haze episode like this occurred at our site frequently during winter. To  
213 understand the influence of wind on nitrate concentration, the wind rose plot is given  
214 in Fig. S1. The prevailing winds at the SORPES station were from northeast and east  
215 during the two-year observation period. Particulate nitrate tended to accumulate or  
216 formed under stagnant condition of low wind speed. The wind from west and east can  
217 lead to a higher nitrate concentration and also other aerosol components (Ding et al.,  
218 2013bc; Shen et al., 2018), which may be associated with air masses from biomass  
219 burning region and the city clusters of YRD.

220 To get an overall picture of nitrate distribution in developed region of costal China,  
221 we reviewed results from available nitrate measurements in three most polluted regions  
222 of North China Plain (NCP), YRD and Pearl River Delta (PRD), and summarized their  
223 concentration, sampling sites and measurement techniques in Fig. 1. Measurements in  
224 summer and winter were separated due to the large seasonal difference of particulate



225 nitrate. Despite that these measurements were from various measurement techniques,  
226 the results still can give us some insights about the differences in spatial and temporal  
227 scales. First, particulate nitrate generally showed the highest concentration in NCP and  
228 followed by YRD and PRD. This was in consistence with the spatial distribution of  $\text{NO}_2$   
229 – a major gas precursor of nitrate. Second, evident seasonal variations can be observed  
230 at all three regions with much higher concentrations in winter. Third, there was an  
231 overall increase trend of particulate nitrate in NCP and YRD in the past decade,  
232 especially that during summertime. Nevertheless, particulate nitrate in PRD revealed  
233 an overall decreasing trend. Compared to these previous studies, the nitrate  
234 concentration during summertime at SORPES station was lower than that in NCP, but  
235 higher than that in YRD and PRD cities. In terms of wintertime, nitrate concentration  
236 at SORPES station was slightly lower than that in NCP, comparable with that in YRD,  
237 and higher than that in PRD.

238 Fig. 2 illustrates the occurrence frequency of the loading of particle matter in  
239 different concentration range, and the changes of nitrate proportion along with the PM  
240 loading. Noting that the PM loading here was indicated by the mass of total WSI. The  
241 highest frequency of WSI concentrations occurred in a range of  $20\text{--}40\ \mu\text{g m}^{-3}$ , and  
242 gradually decreased with the increasing of concentration. Heavy PM pollution with  
243 WSI concentrations higher than  $100\ \mu\text{g m}^{-3}$  occurred during more than 5% of the time  
244 during this study. The contribution of nitrate to total WSI increased with the PM loading,  
245 ranging from  $\sim 25\%$  with WSI concentration lower than  $20\ \mu\text{g m}^{-3}$  to  $\sim 40\%$  when WSI  
246 was higher than  $140\ \mu\text{g m}^{-3}$ . These results suggested that nitrate was a major driver of  
247 haze episodes with high PM peaks in this region.

248 Fig. 3a shows the scatter plot of particulate nitrate and total WSI. They overall  
249 correlated to each other with correlation coefficient  $R = 0.92$ , and nitrate accounted for  
250 32% of the total WSI. Air temperature greatly affected the contribution of nitrate to total  
251 WSI. Its proportion can be up to 58% at around  $0\ ^\circ\text{C}$  and only 3% at the temperature  
252 higher than  $30\ ^\circ\text{C}$ , indicating an important role of thermodynamic equilibrium in nitrate



253 concentration. We further investigated the neutralization extent of sulfate and nitrate by  
254 ammonium (Fig. 3b). Ammonium was overall enough to neutralize both sulfate and  
255 nitrate, suggesting that the particulate nitrate mostly existed as ammonium nitrate at  
256 SORPES station, contrasts with some ammonia poor regions, where excess  $\text{NH}_4^+$   
257 defined as  $([\text{NH}_4^+]/[\text{SO}_4^{2-}]-1.5) \times [\text{SO}_4^{2-}]$  played an important role in the formation of  
258 particulate nitrate (Pathak and Wu, 2009). Evident seasonal difference can be observed  
259 for the molar ratio of ammonium to the sum of sulfate and nitrate. In spring and early  
260 summer, a fraction of the particulate nitrate is present in the forms of  $\text{Ca}(\text{NO}_3)_2$  and  
261  $\text{KNO}_3$ ; while in winter, considerable chloride would consume some ammonium to form  
262  $\text{NH}_4\text{Cl}$ .

263 3.2. Characteristics of fine particular nitrate in different seasons.

264 3.2.1. Seasonal pattern and its main causes

265 Fig. 4 shows the composite seasonal pattern of  $\text{NO}_x$ ,  $\text{PM}_{2.5}$  nitrate, sulfate and the ratio  
266 of nitrate to sulfate during the 2-year period at SORPES station. Similar to the previous  
267 studies (Griffith et al., 2015), a typical seasonal variation was observed for particulate  
268 nitrate (and its ratio to sulfate, i.e.  $\text{NO}_3^-/\text{SO}_4^{2-}$ ), with a maximum value of  $23.7 \mu\text{g m}^{-3}$   
269 (140%) in January, and a minimum of  $8.4 \mu\text{g m}^{-3}$  (66%) in August and September.  
270 Particulate sulfate revealed a bimodal pattern with high concentrations occurred in  
271 January and June, respectively. The low value of particulate nitrate during summer can  
272 be generally explained by the higher temperature, higher and unstable boundary layer  
273 and relative clean air masses induced by summer monsoon (Ding et al., 2013c) despite  
274 of the increased photochemical formation. In opposite, the high values during winter  
275 were generally due to the lower temperature, lower and stable boundary layer and  
276 relative stronger continental outflow from the North China where anthropogenic  
277 emission was relatively high due to heating in winter (Ding et al., 2013c). Different  
278 chemical processes that affects nitrate concentrations between summer and winter will  
279 be discussed later.  $\text{NO}_x$ , the major precursor, tracked the changes of particulate nitrate  
280 commendably, except for that during February and June. In addition, a secondary peak



281 of particulate nitrate can be observed during June, which can be explained as the  
282 influence from agricultural burning in eastern China (Ding et al., 2013bc; Xie et al.,  
283 2015; Shen et al., 2018). The concentrations of Potassium, a biomass burning tracer in  
284 this region (Ding et al., 2013b; Xie et al., 2015), clearly showed a consistent peak (Fig.  
285 S2) with both particulate nitrate and sulfate, as well as the discrepancy of  $\text{NO}_x$  and  
286 nitrate concentrations (Ding et al., 2013b; Xie et al., 2015; Nie et al., 2015) shown in  
287 Fig. 4. While in February, the nitrate concentration didn't show concurrent decrease in  
288  $\text{NO}_x$  during the Chinese Spring Festival (Ding et al., 2013c). It might suggest that local  
289 emissions were not the major contributor to the particulate nitrate observed in winter.

290 Thermodynamics is an important factor influencing the formation and partitioning  
291 of nitrate. Here, we calculated the dissociation constant ( $K_p$ ) to evaluate whether  
292 ambient conditions favored the formation of aerosol  $\text{NH}_4\text{NO}_3$ . The hourly  $\text{NH}_3$   
293 concentration was measured by MARGA and  $\text{HNO}_3$  was calculated by ISORROPIA II.  
294 As showed in Fig. 5, the product of  $[\text{NH}_3(\text{g})]$  and  $[\text{HNO}_3(\text{g})]$  for most ambient samples  
295 during summer was lower than the calculated  $K_p$ , suggesting ambient condition during  
296 summer at SORPES favored to evaporate ammonium nitrate to gas phase  $\text{NH}_3$  and  
297  $\text{HNO}_3$  (Fig. 5a). In contrast, ambient conditions during winter favored to form  
298 particulate nitrate in most cases (Fig. 5b). This study also shows the similar result with  
299 the study reported in the Kanpur, India (Sharma et al., 2007).

### 300 3.2.2. Diurnal cycles during summer and winter

301 In Fig. 6, we show the averaged diurnal variations of particulate nitrate, nitrogen  
302 dioxide, nitric acid, equilibrium constant ( $K$ ), air temperature and RH during summer  
303 and winter during the two years. Nitric acid was calculated by ISORROPIA II. In  
304 summer (Fig. 6a), the fine particulate nitrate showed a typical diurnal cycle that the  
305 maximum concentration occurred at 7:00 with the average concentration of  $16.5 \mu\text{g m}^{-3}$   
306 and minimum value at 14:00 ( $7.2 \mu\text{g m}^{-3}$ ). This summertime diurnal pattern of nitrate  
307 is very similar with the findings in Shandong (Wen et al., 2015) and New York (Sun et  
308 al., 2011). However, it is quite different from the findings in Hong Kong (Griffith et al.,



309 2015), where nitrate concentration peaks in the daytime in summer. Ambient  
310 temperature and the development of boundary layer are the major drivers to the  
311 observed diurnal variation of particulate nitrate, and high temperature and high  
312 boundary layer during daytime prefer to evaporate and dilute the particulate nitrate  
313 (Zhang et al., 2015a; Ding et al., 2016). Nitric acid, which accounted for 20% of the  
314 total nitrate  $[\text{NO}_3^- + \text{HNO}_3]$ , revealed its high concentration (around 2 ppb) in the  
315 noontime (12:00-15:00).  $\text{NO}_2$ , the precursor of nitrate, showed a peak concentration of  
316 18.2 ppb at 21:00, and remained at a high level during the whole night. Equilibrium  
317 constant,  $K$ , was calculated to understand the influence of gas-to-particulate partitioning  
318 on the observed diurnal variation of particulate nitrate (Sun et al., 2011). As showed in  
319 Fig. 6a,  $K$  was highly correlated to particulate nitrate, suggesting the thermodynamic is  
320 the major factor influencing the diurnal variation of particle nitrate during summer.

321 In winter, the diurnal variation is small with a considerable peak appeared at  
322 around 10:00 AM. Compared to that in summer,  $K$  showed similar diurnal variation,  
323 but not correlated to particulate nitrate, indicating factors other than the control of  
324 temperature. The observed peak at late morning was probably due to downward  
325 mixing from the residual layer where particulate nitrate was formed aloft during the  
326 night and brought to the surface after sunrise following the breakup of the boundary  
327 layer (Brown and Dube., 2007; Young et al., 2016; Pusede et al., 2016). Direct vertical  
328 observations are needed to further investigate this issue.

329 To further investigate factors influencing the nitrate behaviors other than  
330 thermodynamics, ISORROPIA II was used to simulate the diurnal variation of nitrate.  
331 Hourly concentrations of all species (both gas and aerosol phase species) at 00:00 were  
332 used as the initial value of each specific day. Hourly data of temperature and relative  
333 humidity were used as the input data to constrain the model. The ISORROPIA II model  
334 was set as forward mode and metastable phase state. The calculated diurnal variations  
335 were showed in Fig. 7 together with the observed results.

336 The differences between the calculation and the observation could be attributed to:



337 (1) the development of boundary layer, (2) the dry deposition of nitric acid, and (3)  
338 chemical processes, which has not been considered yet in the model. As shown in Fig.  
339 7a, the overall diurnal pattern of nitrate in summer is well captured by the model except  
340 for 3 periods. The differences after midnight are likely caused by the effect of boundary  
341 layer height and some chemical processes. Faster increase of model nitrate after 18:00  
342 was attributed to lack of dry deposition of nitric acid in the model. During noontime the  
343 observed nitrate concentration was expected to be lower than the calculated value  
344 because of the development of boundary layer and stronger dry deposition of nitric acid  
345 associated with stronger turbulence mixing, which were ignored in the model. However,  
346 in contrast, the observation was considerably higher than the calculated value. It  
347 indicates a strong production of nitrate via photochemical processes in summer. Fig.7c  
348 shows that the difference between calculated and observed nitrate concentration was in  
349 good correlation with the production of  $\text{NO}_2$  and solar radiation, a proxy for production  
350 rate of nitric acid (Zhang et al., 2005; Young et al., 2016), further suggesting that photo-  
351 oxidation of  $\text{NO}_2$  is an important source of nitrate during summer, even though the  
352 thermodynamic equilibrium is the dominate factor controlling the diurnal cycle. Wen et  
353 al. (2018) also demonstrated that photochemical production of nitric acid is a major  
354 contributor to daytime nitrate increase during summer in North China Plain. In winter,  
355 the influence of thermodynamics is small since the low temperature. The peaks in the  
356 morning may be caused by the mixing down of a residual layer enriched of nitrate as  
357 mentioned above, while the decline during the afternoon are supposed to be the result  
358 of dilution associated with the boundary layer development.

### 359 3.2.3 Influence of air masses transport

360 Meteorological processes play a key role in air masses long-range transport and local  
361 accumulation (Ding et al., 2013ac; Zhang et al., 2016; Ding et al, 2016b). In order to  
362 investigate the influence of air masses transport on nitrate concentrations, Lagrangian  
363 dispersion modeling was conducted for the sampling days with the highest and lowest  
364 25<sup>th</sup> percent nitrate concentration in summer and winter, respectively (Ding et al.,



365 2013a). Fig. 8 shows the retroplumes, i.e. footprint at an altitude of 100 m, of the  
366 selected days during summer and winter, respectively. In summer, high concentrations  
367 of nitrate tended to accompany with the air masses from west of Nanjing (mostly Anhui  
368 province) and Yangtze River Delta (Suzhou-Shanghai city clusters and North Zhejiang  
369 province) (Fig. 8a). YRD is a high  $\text{NO}_x$  emission region (Fig. 1), air masses from which  
370 could bring high concentration of  $\text{NO}_x$  to enhance the nitrate concentration at SORPES  
371 station. Biomass burning is the possible cause of the high nitrate loading with air mass  
372 from the west of Nanjing (Fig. S3). In winter, regional transport from northern China  
373 played an important role in enhancing nitrate concentrations. As shown in Fig. 8c, a  
374 large part of air masses for the highest 25% sampling days was from North China Plain,  
375 which has the strongest  $\text{NO}_x$  emission in China (Fig. 1). It should be noted here that the  
376 longer lifetime of particulate nitrate during winter might be the main cause to promote  
377 the contribution of regional transport to the observed nitrate at SORPES. In contrast,  
378 the lowest 25% sampling days during winter tended to be accompanied with the air  
379 mass from Nanjing local and marine areas.

### 380 3.3. Contribution of $\text{N}_2\text{O}_5$ hydrolysis to nitrate episodes

381 Similar to findings from previous studies (Zhang et al., 2015c), nitrate was found to  
382 increase significantly during this study and become the largest contributor of  $\text{PM}_{2.5}$   
383 during the haze episodes (Fig. 2). Generally, these pollution episodes mainly occurred  
384 in winter (Fig. 3a and Fig. 4), during which the photochemical production of nitric acid  
385 should be weak.  $\text{N}_2\text{O}_5$  hydrolysis was thus proposed to be a potential important  
386 formation pathway. Here we investigated the nitrate episodes in detail and discussed  
387 their relationship to the  $\text{N}_2\text{O}_5$  hydrolysis during the nights before.

388 In Fig. 9, we show a typical case of nitrate episodes from 30 November to 2  
389 December, 2015. Fast nitrate formation was observed, which was likely caused by  
390 hydrolysis of  $\text{N}_2\text{O}_5$ . Nitrate increased significantly from  $20.3 \mu\text{g m}^{-3}$  at 18:00 of 30  
391 November to  $63 \mu\text{g m}^{-3}$  at 6:00 of 1 December, 2015. The ratio of nitrate to  $\text{PM}_{2.5}$  also  
392 exhibited a large increase from 25% at 18:00 to 38% at 06:00. In contrast, other  $\text{PM}_{2.5}$



393 components, e.g. sulfate and black carbon, showed only slight increases. High  
394 concentration of  $\text{NO}_2$ , considerable level of  $\text{O}_3$  and extremely low concentration of  $\text{NO}$   
395 provided a favorable condition towards forming  $\text{NO}_3$  and  $\text{N}_2\text{O}_5$  (Brown et al., 2003).  
396 The meteorological conditions during these 12 hours were stable with low wind speed  
397 and high relative humidity, which, combined with the relatively high concentration of  
398  $\text{PM}$ , would promote the hydrolysis of  $\text{N}_2\text{O}_5$  (Riemer, 2003).

399  $\text{N}_2\text{O}_5$  concentrations were calculated by using steady-state expressions (Osthoff et  
400 al., 2006; Wang et al., 2014; Wen et al., 2015), and the result was shown in Fig. 9. The  
401 calculated  $\text{N}_2\text{O}_5$  exhibited a much higher concentration during the night of 30  
402 November compared to the days before and after. Particulate nitrate formed from  $\text{N}_2\text{O}_5$   
403 hydrolysis was then computed during the 12-hour period. Nitrate concentration at 18:00  
404 of 30 November, 2015 ( $20 \mu\text{g m}^{-3}$ ) was selected as the initial value, and  $31 \mu\text{g m}^{-3}$  of  
405 particulate nitrate was produced in the following 12 hours, suggesting that approximate  
406 80% of increased particulate nitrate can be attributed to the hydrolysis of  $\text{N}_2\text{O}_5$  in this  
407 case.

408 To further understand the contribution of  $\text{N}_2\text{O}_5$  hydrolysis, sampling days with  
409 daily-averaged nitrate concentration exceeding the mean plus twice the standard  
410 deviation were selected as the nitrate episode days. In total, 18 episode-days were  
411 selected during the 2-year measurement, with 16 days in winter and the other 2 days in  
412 biomass burning season. In Fig. 10, we presented the averaged diurnal pattern of  
413 particulate nitrate and its related parameters on the 18 selected episode and pre-episode  
414 days. For the episode days, particulate nitrate revealed a similar diurnal pattern as that  
415 of the whole winter (Fig. 6). Nitrate maintained a high concentration during the whole  
416 day with a small peak around 10:00 in the morning. However, for the pre-episode days,  
417 a clear build-up of nitrate can be observed, especially during the nighttime from 17:00  
418 of the pre-episode days to 1:00 of the episode days (as marked in Fig. 10). The average  
419 increment of ammonium nitrate has exceeded  $24 \mu\text{g m}^{-3}$  during this 9-hour period of  
420 the pre-episode nighttime. The total WIS also increased during this period, which was



421 mostly attributed to ammonium nitrate (almost 90%) and resulted in an evident increase  
422 of the ratio of nitrate to total WSI. Compared to nitrate, black carbon, a tracer of primary  
423 emissions, showed little change during the pre-episode day. The retroplume showed in  
424 Fig. S4 suggested that the air masses arrived at the SORPES station on the pre-episode  
425 and episode days were almost the same. These results suggest that secondary formation  
426 other than accumulation was the major contributor to the observed increase of  
427 particulate nitrate.

428 Since the observed nitrate formation mostly occurred during the nighttime of pre-  
429 episode days when the photochemical production of nitric acid would be largely  
430 suppressed,  $N_2O_5$  hydrolysis is thus believed to be the major contributor. As showed in  
431 Fig. 10, compared to those during episode days,  $NO_2$  concentration was comparable,  
432 but  $O_3$  concentration was higher during pre-episode days. This resulted in a higher  
433 production rate of  $N_2O_5$  proxy ( $NO_2 \times O_3$ ) in pre-episode days, and favored formation of  
434 nitrate from the hydrolysis of  $N_2O_5$ . We further calculated the contribution of  $N_2O_5$   
435 hydrolysis to nitrate formation during the periods from 17:00 to 23:00 of each pre-  
436 episode day (excluding 2 windy days). A good correlation ( $R=0.8$ ) was observed  
437 between the calculated nitrate and observed nitrate (Fig. 11), with the slope of 0.8,  
438 indicating most of the observed nitrate formation during nitrate episodes were attributed  
439 to the hydrolysis of  $N_2O_5$ .

#### 440 4. Summary and Conclusion

441 Online measurements of fine particulate nitrate along with trace gases and  $PM_{2.5}$  mass  
442 concentrations were conducted for two years from March 2014 to February 2016 using  
443 a MARGA at SORPES station, a rural receptor site in the Yangtze River Delta, eastern  
444 China. Hourly nitrate concentration varied from  $0.5 \mu g m^{-3}$  to  $92.6 \mu g m^{-3}$ , with an  
445 averaged value of  $15.8 \mu g m^{-3}$ , which was generally higher than the measurement at the  
446 sites in YRD and PRD, but lower than that at the sites in North China Plain. The  
447 contribution of nitrate to total WSI increased from 25% with WSI concentration lower  
448 than  $20 \mu g m^{-3}$ , to 40% when WSI was higher than  $140 \mu g m^{-3}$ , suggesting a major



449 driver of nitrate to the aerosol pollution in YRD.  $\text{NH}_3$  is enough to neutralize the acidic  
450 compounds of aerosol, ammonium nitrate was thus the predominate form of the  
451 observed particulate nitrate. A clear seasonal variation of nitrate was observed with  
452 peak value in January and December and lowest value in August and September.  
453 Biomass burning plumes contributed the nitrate concentration evidently and resulted in  
454 a secondary peak during June. In summer, thermodynamic equilibrium was the major  
455 factor influencing the diurnal variation of nitrate, and resulted in a much lower  
456 concentration at noontime. Nevertheless, the observed nitrate at noontime was  
457 considerably higher than the value predicted by ISORROPIA II model, indicating a  
458 strong production of nitrate by the photo-oxidation of  $\text{NO}_2$ . Air masses from YRD and  
459 biomass burning region were corresponded to the high nitrate concentrations during  
460 summer. In winter, the diurnal variation of nitrate was weak. Regional transport from  
461 North China Plain contributed largely to the observed high nitrate concentrations.

462 Nitrate episodes, defined as daily-averaged concentration exceeding the mean  
463 value plus twice the standard deviation, were further investigated to understand the  
464 chemical processes towards forming particulate nitrate and their contribution to the  
465 pollution episodes. A clear build-up of nitrate can be observed during the pre-episode  
466 night, and dominated the increase of total WSI.  $\text{N}_2\text{O}_5$  hydrolysis was demonstrated to  
467 contribute 80% of the observed nitrate formation, suggesting its critical role in an  
468 aerosol pollution episode. In view of the significant emission of  $\text{NO}$ , which is the main  
469 sink of  $\text{N}_2\text{O}_5$  during night, stronger production of  $\text{N}_2\text{O}_5$  is expected at the upper  
470 boundary layer, e.g. residual layer, and contribute to the nitrate formation in the entire  
471 boundary layer. In summary, our study provides evidence that particulate nitrate  
472 especially that formed from  $\text{N}_2\text{O}_5$  hydrolysis is a crucial contributor to the aerosol  
473 pollution episodes in eastern China.

474 *Data availability.* The GDAS data used in the HYSPLIT calculation can be acquired  
475 from <ftp://arlftp.arlhq.noaa.gov/pub/archives/gdas1>. Measurement data at SORPES,



476 including aerosol data and relevant trace gases as well as meteorological data, are  
477 available upon request from the corresponding author before the SORPES database is  
478 opened publicly.

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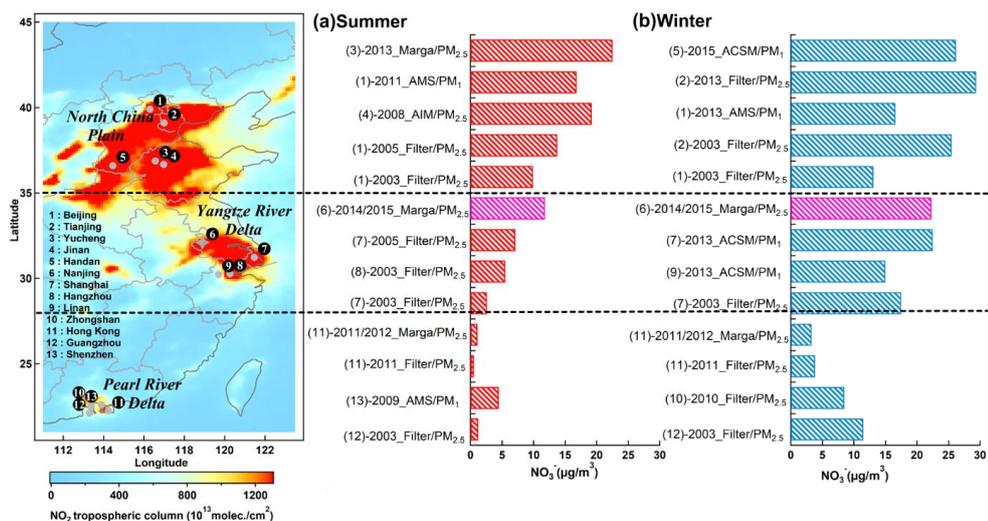
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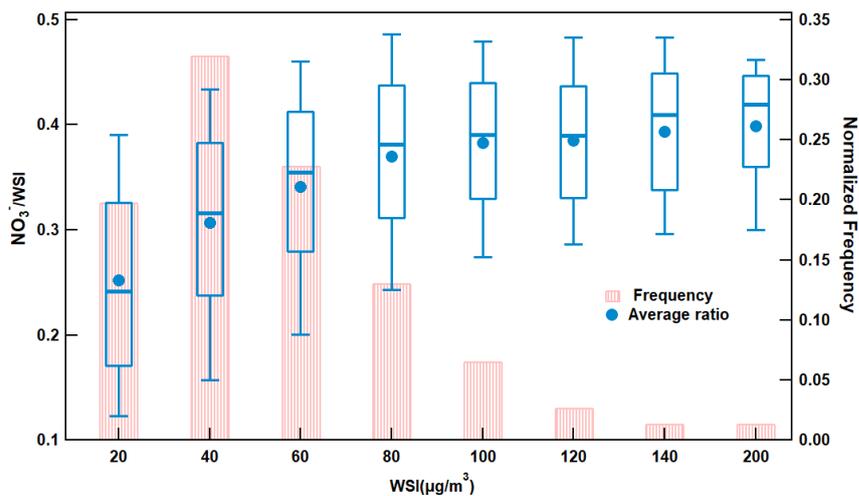
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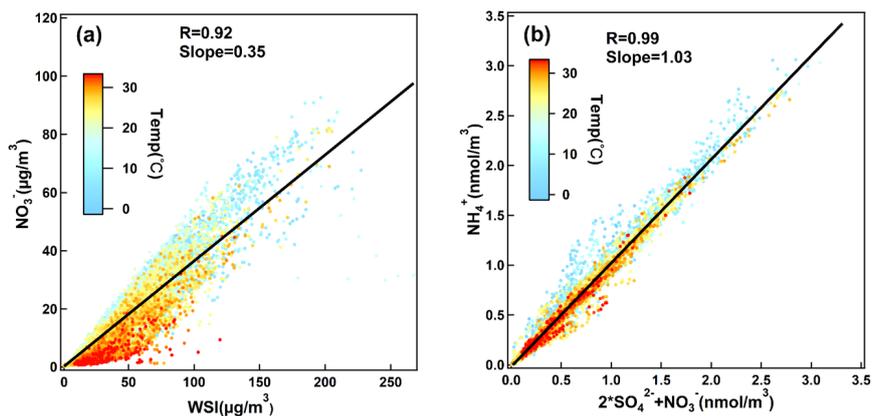
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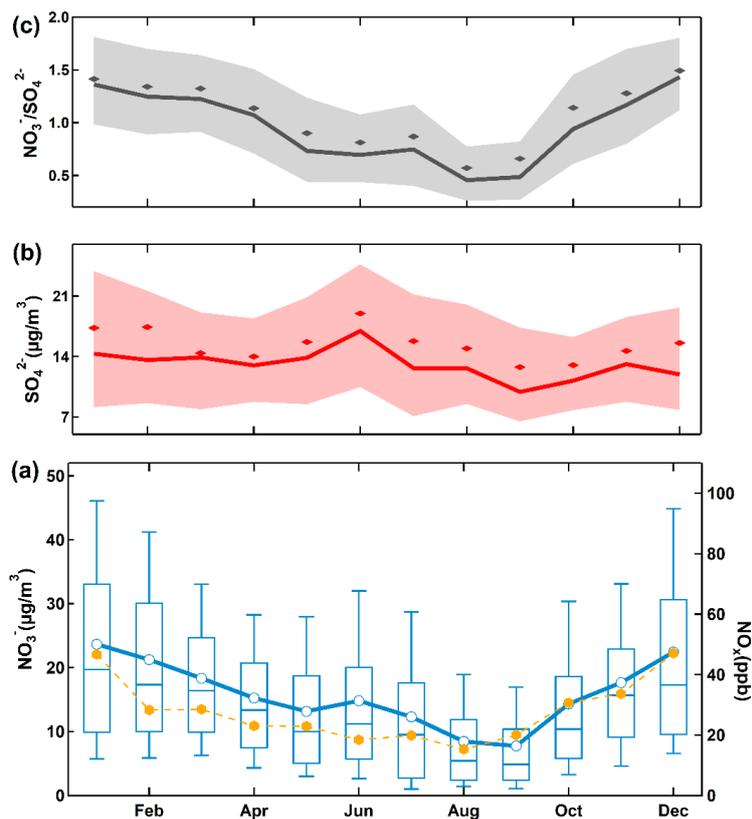
**Figure 1** Average mass concentrations of particulate nitrate at different sampling sites in **(a)** summer and **(b)** winter. The left panel shows the map color-coded by 2-years (2014-2015) averaged tropospheric NO<sub>2</sub> from OMI satellite (<http://www.temis.nl/airpollution/no2.html>). The pink bars are for this study.



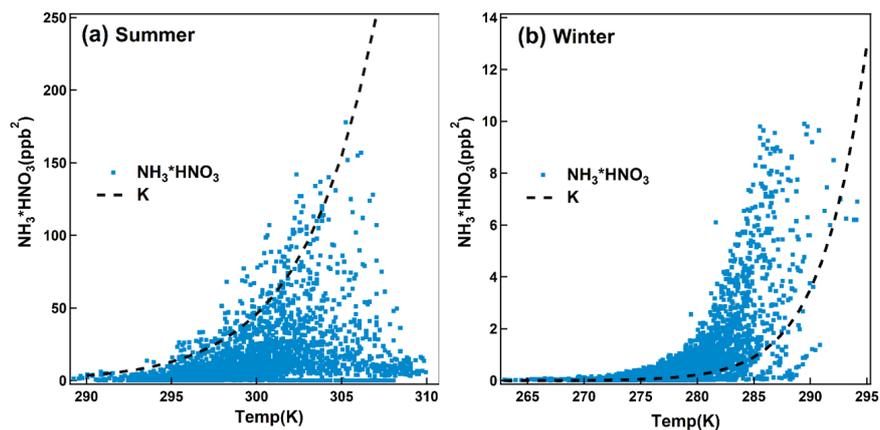
**Figure 2** Average proportion of nitrate and normalized frequency of occurrence at different mass concentration bins of water soluble ions at SORPES. For the ratio, box boundaries represent the interquartile range, bars represent 5%-95% percentile range, and horizontal lines represent the median value.



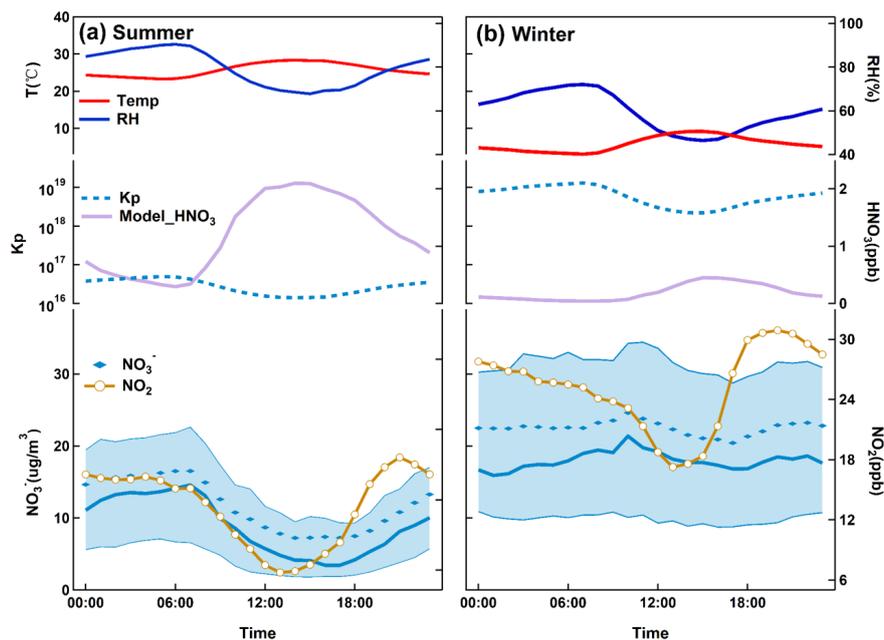
**Figure 3** Scatter plots of (a) nitrate vs. total WSI color coded by air temperature, (b) molar concentrations of ammonium with nitrate molar concentrations plus two times of sulfate molar concentrations.



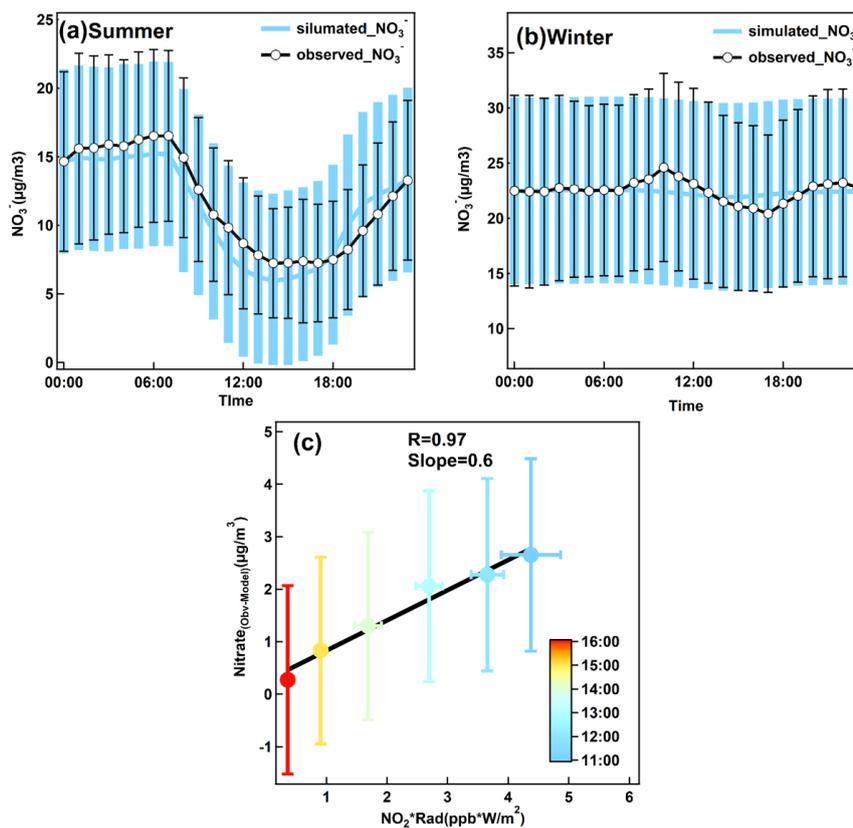
**Figure 4** Monthly averaged nitrate (blue), sulfate (red),  $\text{NO}_x$  (orange) mass concentrations and nitrate to sulfate ratio (grey) measured at SORPES station during March 2014 to February 2016. For nitrate to sulfate ratio (a) and sulfate (b), bold solid lines are the median values, shade areas represent percentiles of 75% and 25%, and diamonds represent the mean values. For nitrate (c), box boundaries represent the interquartile range, bars represent 5%-95% percentile range, horizontal lines represent the median value, and crosses represent mean values.



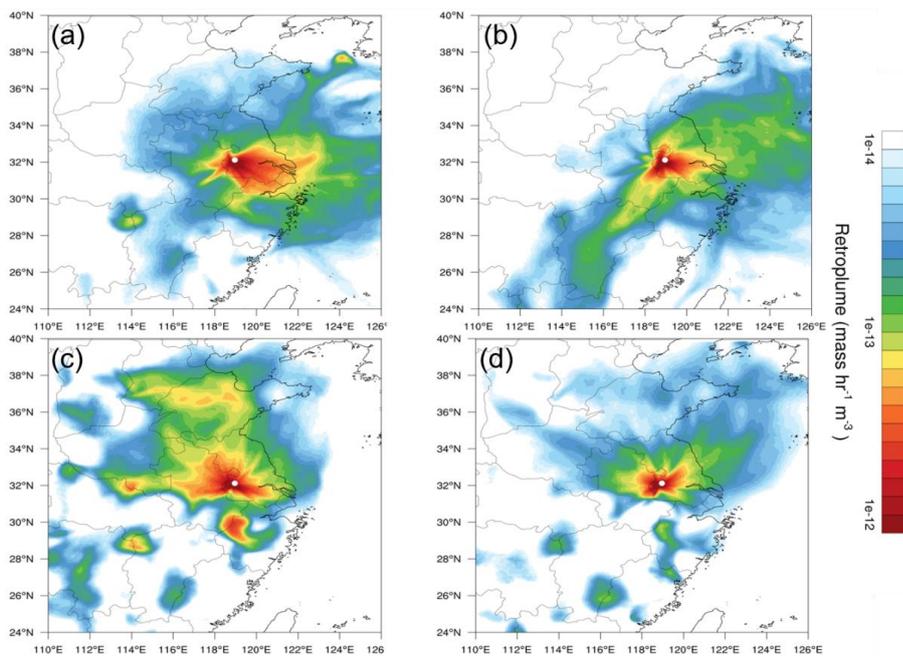
**Figure 5** Scatter plot of product of  $[\text{NH}_3(\text{g})] \cdot [\text{HNO}_3(\text{g})]$  with temperature together with the K parameter calculated by temperature for (a) summer and (b) winter.



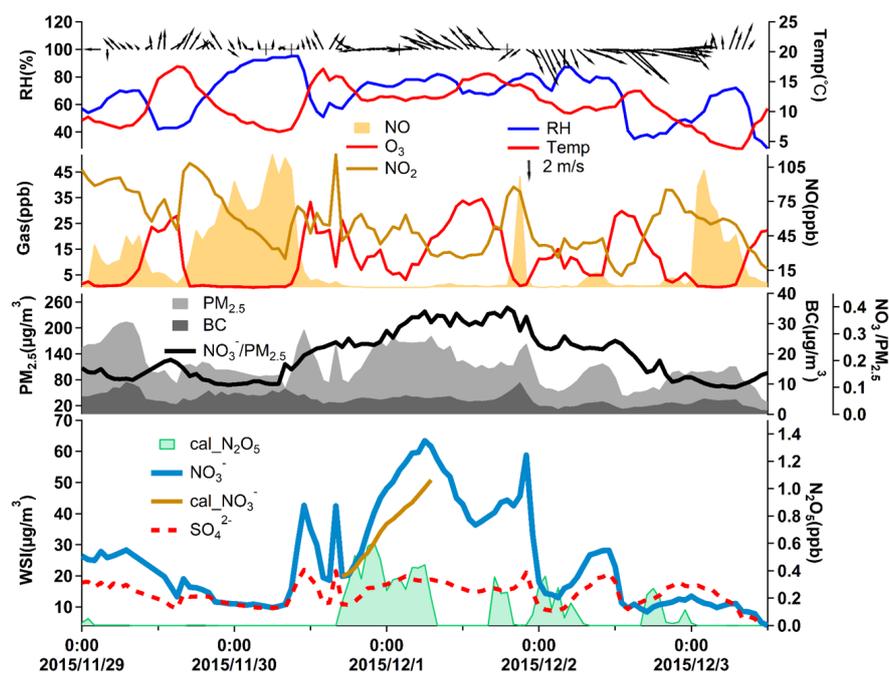
**Figure 6** Diurnal variation of particulate nitrate in **(a)** summer and **(b)** winter. For nitrate, bold solid lines are the median values, shaded areas represent percentiles of 75% and 25% and diamonds represent mean values. Diurnal averages of  $\text{NO}_2$  and modeled nitric acid mass concentrations are also provided with temperature and RH.



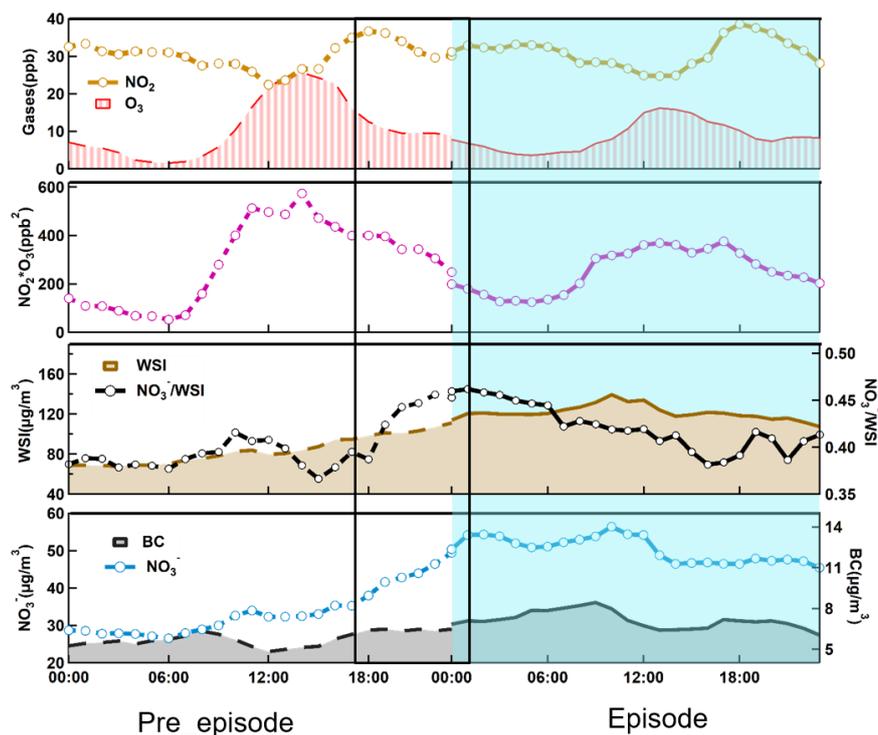
**Figure 7** Modeled nitrate diurnal variations in (a) summer and (b) winter, together with the observed nitrate concentrations. Error bars provided are the standard deviation of the mean at each hourly interval. (c) Scatter plot of the difference between model and observed nitrate average mass concentrations with the product of NO<sub>2</sub> and radiation color coded by the hour of day for the samples.



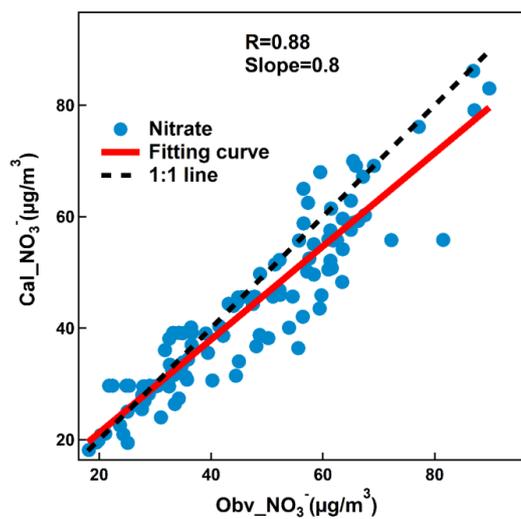
**Figure 8** The averaged retroplumes (i.e., 100 m footprint) of the selected events: **(a)** Top 25% nitrate concentrations in summer, **(b)** Bottom 25% nitrate concentrations in summer, **(c)** Top 25% nitrate concentrations in winter, and **(d)** Bottom 25% nitrate concentrations in winter.



**Figure 9** Time series of meteorological data and the concentrations of trace gases related to nitrate formation during 29 November to 3 December, 2015. Cal\_NO<sub>3</sub><sup>-</sup> represents the nitrate concentrations calculated from the hydrolysis of N<sub>2</sub>O<sub>5</sub>.



**Figure 10** Diurnal variations of particulate nitrate, black carbon, the total water soluble ions, nitrate to WSI ratio, product of  $\text{NO}_2$  and  $\text{O}_3$ ,  $\text{NO}_2$ , and  $\text{O}_3$  averaged for nitrate episode days with exceedances of one mean plus two standard deviations. The left side shows the pre-episode days and the right side shows the episode days during the winter of entire two years period. The solid line box corresponds to the rapid growth of nitrate at night.



**Figure 11** Scatter plot of calculated nitrate concentrations and observed nitrate concentrations from 17:00 to 23:00 of each episode.

Table 1 major gas phase and heterogeneous reactions involved NO<sub>3</sub> and N<sub>2</sub>O<sub>5</sub>

Reaction	Rate constant
$\text{NO}_2 + \text{O}_3 \rightarrow \text{NO}_3 + \text{O}_2$	$k_1$
$\text{NO}_3 + \text{NO}_2 \leftrightarrow \text{N}_2\text{O}_5$	$k_{eq}$
$\text{NO}_3 + \text{NO} \leftrightarrow \text{NO}_2 + \text{NO}_2$	$k_3$
$\text{NO}_3 \rightarrow \text{NO} + \text{O}_2$	$j_4$
$\text{NO}_3 \rightarrow \text{NO} + \text{O}_2$	$j_5$
$\text{NO}_3 \xrightarrow{\text{voc}} \text{products}$	$k_6 = \sum (k_{\text{voci}} \cdot [\text{voc}]_i)$
$\text{NO}_3 \xrightarrow{\text{Heterogeneous}} \text{products}$	$k_7 = 0.25 \cdot i \cdot C_{\text{NO}_3} \cdot \gamma_{\text{NO}_3} \cdot S_{\text{aerosol}}$
$\text{N}_2\text{O}_5 \xrightarrow{\text{Heterogeneous}} \text{products}$	$k_8 = 0.25 \cdot C_{\text{N}_2\text{O}_5} \cdot \gamma_{\text{N}_2\text{O}_5} \cdot S_{\text{aerosol}}$
$\text{N}_2\text{O}_5 \xrightarrow{\text{Homogeneous}} 2\text{HNO}_3$	$k_9 = k_{\text{I}} [\text{H}_2\text{O}] + k_{\text{II}} [\text{H}_2\text{O}]^2$