

1 **Two years online measurement of fine particulate nitrate in western**  
2 **Yangtze River Delta: Influences of thermodynamics and N<sub>2</sub>O<sub>5</sub> hydrolysis**

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11 **Abstract.**

12 Particulate nitrate contributes a large fraction of secondary aerosols. Despite  
13 understanding of its important role in regional air quality and global climate, long-term  
14 continuous measurements are rather limited in China. In this study, we conducted online  
15 measurement of PM<sub>2.5</sub> nitrate for two years from March 2014 to February 2016 using  
16 the Monitor for Aerosols and Gases in ambient Air (MARGA) in the western Yangtze  
17 River Delta (YRD), eastern China, and investigate the main factors that influenced its  
18 temporal variations and formation pathways. Compared to other sites in China, an  
19 overall high concentration of particulate nitrate was observed with a mean value of 15.8  
20  $\mu\text{g m}^{-3}$  (0.5 to 92.6  $\mu\text{g m}^{-3}$ ). Nitrate on average accounted for 32% of the total mass of  
21 water-soluble ions and the proportion increased with PM loading, indicating that nitrate  
22 is a major driver of haze pollution episodes in this region. Sufficient ammonia drove  
23 most nitrate into the particle phase in the form of ammonium nitrate. A typical seasonal  
24 cycle of nitrate was observed with the concentrations in winter on average two times  
25 higher than those in summer mainly due to different meteorological conditions. In  
26 summer, the diurnal variation of particulate nitrate was determined by the  
27 thermodynamic equilibrium, resulting in a much lower concentration during daytime  
28 despite of a considerable photochemical production. Air masses from polluted YRD

29 and biomass burning region contributed to the high nitrate concentration during summer.  
30 In winter, particulate nitrate didn't reveal an evident diurnal variation. Regional  
31 transport from northern China played an important role in enhancing nitrate  
32 concentration. Eighteen nitrate episodes were selected to understand the processes that  
33 drive the formation of high concentration of nitrate. Rapid nitrate formation was  
34 observed during the pre-episode (the day before nitrate episode day) nights, and  
35 dominated the increase of total water-soluble ions. Calculated nitrate from  $N_2O_5$   
36 hydrolysis was highly correlated to and accounted for 80 percent of the observed nitrate,  
37 suggesting that  $N_2O_5$  hydrolysis was a major contributor to the nitrate episodes. Our  
38 results suggested that rapid formation of nitrate could be a main cause for extreme  
39 aerosol pollution events in YRD during winter, and illustrated the urgent needs to  
40 control the  $NO_x$  emission.

## 41 **1. Introduction**

42 Particulate nitrate ( $NO_3^-$ ), as a major aerosol component in the atmosphere, reduces  
43 atmospheric visibility (Charlson and Heintzenberg, 1995), influences human health,  
44 alters radiative forcing and hence influences regional even global climate (IPCC, 2013).  
45 Compared to the sulfate, nitrate has a larger scattering albedo under low RH conditions  
46 that cause a stronger influence on visibility (Lei and Wuebbles, 2013). High  
47 concentration of particulate nitrate had been demonstrated to be one of the major  
48 reasons for the frequent occurrence of haze episodes in China (Wang and Zhang, 2009;  
49 Wen et al., 2015; Wang et al., 2017). In recent decades, the Chinese government started  
50 to control emissions of air pollutants with special effort on the  $SO_2$  reduction. This  
51 resulted in a remarkable decrease of ambient  $SO_2$  and sulfate concentrations after 2006  
52 (van der A et al., 2017; Wang et al., 2017). However, particulate nitrate, as well as its  
53 proportion in PM, showed increasing trends due to the strong emission of nitrogen  
54 oxides ( $NO_x$ ) (Lei and Wuebbles, 2013; Yang et al., 2017).

55 Particulate nitrate can be formed from multiple pathways. Gas phase reaction of  
56  $NO_2$  and OH radical is one major pathway to form nitric acid ( $HNO_3$ ) (Calvert and

57 Stockwell, 1983), which subsequently reacts with ammonia ( $\text{NH}_3$ ) to produce  
58 ammonium nitrate ( $\text{NH}_4\text{NO}_3$ ). As typical photochemical processes, these reactions  
59 dominate daytime nitrate formation, and have been widely investigated in both field  
60 and modelling studies (Sharma et al., 2007; Petetin et al., 2016). Heterogeneous uptake  
61 of the photochemical formed nitric acid by alkali compounds, e.g. dust and sea salt  
62 particles, is also a considerable pathway to form nitrate in some regions (Bian et al.,  
63 2014). During nighttime, the hydrolysis of dinitrogen pentoxide ( $\text{N}_2\text{O}_5$ ) is believed to  
64 be the dominate pathway to form particulate nitrate.  $\text{N}_2\text{O}_5$  is an important reactive  
65 nitrogen species in the polluted troposphere (Brown and Dube., 2007; Osthoff et al.,  
66 2006; Li et al., 2017; Brown et al., 2003; Brown and Stutz, 2012; Wang et al., 2016d;  
67 Wang et al., 2017b) and accumulates via the reversible reaction between  $\text{NO}_2$  and  $\text{NO}_3$   
68 radical produced from the reaction of  $\text{NO}_2$  with  $\text{O}_3$ . Due to the rapid photolysis of  $\text{NO}_3$   
69 radical, the contribution of  $\text{N}_2\text{O}_5$  hydrolysis to nitrate concentration during daytime of  
70 sunny day is usually small. While during nighttime,  $\text{N}_2\text{O}_5$  concentration can be up to  
71 ppb level, and form nitric acid by reaction with water vapor, or particulate nitrate  
72 directly by heterogeneous hydrolysis on the wet surface (Wang et al., 2017a; Wen et al.,  
73 2018; Thornton et al., 2003). In China, the pollution episodes with high nitrate  
74 concentrations mostly occurred in winter, during which the photochemical production  
75 of nitrate should be overall weak.  $\text{N}_2\text{O}_5$  hydrolysis thus has the potential to be the  
76 crucial contributor, however there is still a lack of observational evidences.

77 Collecting particulate matter on a filter with subsequent ion chromatography  
78 analysis in laboratories is the conventional method to measure the concentration of  
79 particulate nitrate. The un-denuded filter pack system, which is most-widely used, can  
80 suffer from both positive artifacts by absorbing gas-phase nitric acid, and negative  
81 artifacts by the evaporation of ammonium nitrate (Nie et al., 2010; Pathak and Wu,  
82 2009; Wang et al., 2010). A denuder system can minimize these sampling artifacts by  
83 adding denuders to remove the interfering gases and back-up filters to collect the  
84 evaporated vapors (John et al., 1988). However, the operation of such a denuder system

85 is extremely labor intensive and thus not widely used. In addition, the poor time-  
86 resolution of filter-based measurement can limit our understanding on the formation  
87 and chemical evolution of the particulate nitrate. To overcome these shortcomings,  
88 several continuous and semi-continuous techniques have been developed based on an  
89 online denuded-IC system (e.g. the ambient ion monitor (AIM), the gas and aerosol  
90 collector ion chromatography (GAC-IC), the particle-into-liquid sample ion  
91 chromatography (PILS-IC) and MARGA), as well as mass spectrometry (e.g. AMS).  
92 Pathak et al. (2011) applied an AIM instrument in Beijing and Shanghai for one month  
93 and found that the heterogeneous hydrolysis of  $N_2O_5$  contributed 50%~100% of the  
94 nighttime nitrate formation. Xue et al. (2013) deployed a PILS-IC system in Hong Kong  
95 for less than a month and showed a more active nitrate formation during PM episode  
96 than normal days. Wen et al. (2015) used a MARGA instrument in Yucheng, North  
97 China during summer and emphasized the important roles of  $O_3$  and  $NH_3$  on nitrate  
98 formation. Yang et al. (2017) carried out field observation with ACSM in Beijing for  
99 half a month and pointed out the importance of aerosol nitrate in haze formation.  
100 However, despite of an increasing number of studies using online techniques,  
101 continuous measurements with more than one-year period are still very limited.

102 The Yangtze River Delta (YRD) located in the eastern China, with megacities  
103 including Shanghai, Nanjing and Suzhou etc., has suffered from heavy particulate  
104 matter pollution and photochemical pollution (Ding et al., 2013bc; Wang et al., 2016bc;  
105 Wang et al., 2016a). Previous studies indicated an important role of nitrate in the  
106 pollution episodes (Hua et al., 2015; Du et al., 2011; Yang et al., 2017). For example,  
107 Zhang et al. (2015a) carried out an observation with ACSM in urban Nanjing during  
108 summer and autumn, and found nitrate and organic aerosols dominated the  $PM_{10}$   
109 composition. Shi et al. (2014) used MARGA instrument in Shanghai for months and  
110 found an increasing contribution of nitrate to  $PM_{10}$  mass during pollution periods.  
111 Wang et al. (2016b) reported temporal variation and transport of  $PM_{2.5}$  water soluble  
112 ions, including nitrate, in an urban site in Shanghai based on three-year continuous

113 measurement using MARGA. However, detailed investigation of the possible  
114 mechanisms governing nitrate behaviors during haze pollution is still rare.

115 In this study, we present a 2-year continuous measurement of particulate nitrate  
116 using MARGA at a rural site in Nanjing, a megacity in the western YRD region, with  
117 the target to get a comprehensive understanding of particulate nitrate behaviors and  
118 investigate the processes affecting nitrate in haze episodes. We first conducted general  
119 statistical analysis of particulate nitrate and characterized seasonal variation and diurnal  
120 pattern. A thermodynamic model was then applied to investigate the gas-particle  
121 partition of nitrate. The influence of air masses was also investigated by conducting  
122 backward Lagrangian dispersion modelling. Finally, we selected eighteen nitrate  
123 episodes and investigated the main processes influencing their evolution.

## 124 **2. Methodology**

### 125 2.1. Sample site and instrumentation

126 The SORPES station (118°57'E, 32°07'N) was located on the top of a small hill (40 m  
127 above sea level) in the Xianlin campus of Nanjing University located in the outskirts of  
128 Nanjing, China. The station is an ideal receptor of air masses from the YRD with little  
129 influence of local emissions and urban pollution from Nanjing. Detailed description can  
130 be found in previous studies (Ding et al., 2013c; Ding et al., 2016a).

131 The measurement was conducted from March 2014 to Feb 2016. Hourly  
132 concentrations of water soluble gases of HCl, HNO<sub>3</sub>, HONO, SO<sub>2</sub> and NH<sub>3</sub>, and water-  
133 soluble ions in PM<sub>2.5</sub>, including Cl<sup>-</sup>, NO<sub>3</sub><sup>-</sup>, SO<sub>4</sub><sup>2-</sup>, NH<sub>4</sub><sup>+</sup>, Na<sup>+</sup>, K<sup>+</sup>, Ca<sup>2+</sup>, and Mg<sup>2+</sup>, were  
134 measured with a Monitor for Aerosols and Gases in ambient Air (MARGA, designed  
135 and manufactured by Applikon Analytical B.V., the Netherlands) employed in  
136 connection with a Thermo PM<sub>2.5</sub> cyclone inlet. The sampling system was comprised of  
137 two parts: A Wet Rotating Denuder for gases and a Steam Jet Aerosol Collector for  
138 aerosols, working at an air flow of 1 m<sup>3</sup> h<sup>-1</sup> (ten Brink et al., 2007; Rumsey et al., 2014).  
139 After each hour's collection, the samples were analyzed using ion chromatography. The  
140 instrument was calibrated on an hourly basis using internal standard liquid (bromide

141 lithium), ensuring a stable and reliable ion chromatograph. Concentrations of all aerosol  
142 ions and gases have a precision of  $0.001\mu\text{g m}^{-3}$  (Xie et al., 2015). The  $\text{PM}_{2.5}$  ion dataset  
143 from the MARGA provided more than 15000 hourly samples over the 24 months of  
144 measurements considering points where  $\text{NO}_3^-$ ,  $\text{NH}_4^+$  and  $\text{SO}_4^{2-}$  were all available. Trace  
145 gases (i.e.,  $\text{O}_3$ ,  $\text{SO}_2$ ,  $\text{NO}_x$ ,  $\text{NO}$ ) and  $\text{PM}_{2.5}$  mass concentrations were also measured at  
146 SORPES (Ding et al., 2013c; Nie et al., 2015; Ding et al., 2016a), together with  
147 meteorological data including wind speed/direction, temperature, and relative humidity.

## 148 2.2. Thermodynamic constants and ISORROPIA II

149 Formation of ammonium nitrate involves an equilibrium reaction between the gas phase  
150  $\text{NH}_3$  and  $\text{HNO}_3$ , and particle phase  $\text{NH}_4\text{NO}_3$ . The gas-to-particle partitioning is  
151 temperature dependent, and the equilibrium constant can be calculated as follows (units  
152  $\text{mol}^2\text{kg}^{-2}\text{atm}^{-2}$ ) (Sun et al., 2011; Seinfeld and Pandis, 2006):

$$153 \quad K=k_{298} \exp(a(298/T-1) +b[1+\ln(298/T)-298/T]) \quad (1)$$

$$154 \quad K_{298}=3.5\times 10^{16}(\text{atm}^{-2}), a=75.11, b=-13.5. \quad (2)$$

155 ISORROPIA II (available at <http://isorroopia.eas.gatech.edu/>) is a thermodynamic  
156 model used commonly in inorganic aerosol research (Fountoukis and Nenes, 2007). To  
157 analyze gas-into-particulate pathway for nitrate formation,  $\text{HNO}_3$  was modeled with  
158 ISORROPIA II run in forward model iteratively (Pusede et al., 2016; Fountoukis and  
159 Nenes, 2007). ISORROPIA II was initialized as  $[\text{NO}_3^-+\text{HNO}_3]_{\text{total}} = [\text{NO}_3^-]_{\text{aerosol}}$ .  
160 Calculated  $\text{HNO}_3(\text{g})$  was added back to  $[\text{NO}_3^-+\text{HNO}_3]$ , while we always use  
161  $[\text{NH}_4^+_{\text{aerosol}}+\text{NH}_3(\text{g})]$  as input total ammonium. ISORROPIA II was solved iteratively  
162 until output  $\text{NO}_3^-$  changed by  $< 2\%$  by mass. The phase state was set as metastable. We  
163 assume that gases and aerosol are in equilibrium, that aerosols are homogeneous and  
164 internally mixed, and that unaccounted-for factors do not influence the thermodynamics  
165 of system (Vayenas et al., 2005).

## 166 2.3. Lagrangian Dispersion Modeling

167 To help understand the influence of air masses, backward Lagrangian particulate  
168 dispersion modeling (LPDM) was carried out based on a method developed and

169 evaluated by Ding (Ding et al., 2013a). The LPDM was conducted using the Hybrid  
170 Single-Particulate Lagrangian Integrated Trajectory model developed in the Air  
171 Resource Laboratory (ARL) of the National Oceanic and Atmospheric Administration  
172 using the ARL format Global Data Assimilation System data. The model calculates the  
173 position of particulates by mean wind and a turbulence transport component after they  
174 are released at the source point for a backward simulation. For each hour, 3000  
175 particulates were released at 100 m altitude over the site and were traced backward for  
176 a 3-day period. The hourly position of each particulate was calculated using a 3-D  
177 particulate, i.e., horizontal and vertical, method. The residence time at 100 m altitude,  
178 i.e., foot-print “retroplume”, which represents the distribution of the surface probability  
179 or residence time of the simulated air mass, was used to understand the contribution  
180 from potential source regions (Ding et al., 2013ac; Shen et al., 2018).

#### 181 2.4. Steady-state predictions

182 Based on their short lifetimes, the concentrations of the  $\text{NO}_3$  radical and  $\text{N}_2\text{O}_5$  can be  
183 predicted by steady-state calculations due to lack of measurement data (Osthoff et al.,  
184 2006). The formation and loss of  $\text{N}_2\text{O}_5$  associated with a series of chemical reactions  
185 are listed in Table 1. For the heterogeneous processes, we used 0.004 and 0.03 as the  
186 uptake coefficients of the  $\text{NO}_3$  radical and  $\text{N}_2\text{O}_5$  ( $\gamma\text{NO}_3$  and  $\gamma\text{N}_2\text{O}_5$ ), respectively  
187 (Aldener et al., 2006; Wen et al., 2015; Knopf et al., 2011; Brown et al., 2006). Due to  
188 the lack of measurement during 2014 - 2016, the VOCs data used here was the average  
189 of nighttime value that measured at SORPES station during the wintertime of 2017.  
190 The corresponding rate constant can be found in Master Chemical Mechanism (MCM  
191 version 3.1, <http://mcm.leeds.ac.uk/MCM/>). The uncertainty that caused by the limited  
192 VOCs measurements and the variation of  $\text{N}_2\text{O}_5$  uptake coefficient were estimated and  
193 discussed in the supplement.

### 194 3. Results and discussion

#### 195 3.1. Overall results

196 A MARGA was deployed to continuously measure the eight water-soluble ions (WSI)

197 of PM<sub>2.5</sub> and several gas-phase species from March 2014 to February 2016 with the  
198 time resolution of 1-hour. Nitrate (NO<sub>3</sub><sup>-</sup>), sulfate (SO<sub>4</sub><sup>2-</sup>) and ammonium (NH<sub>4</sub><sup>+</sup>) were  
199 the major components with the two-year averaged concentrations of 15.8 (±13.4), 15.3  
200 (±10.6) and 10.4 (±7.6) μg m<sup>-3</sup>, respectively. In the present study, we focused on nitrate,  
201 and discussed the temporal variation and its association with physicochemical  
202 processes. The concentration of PM<sub>2.5</sub> nitrate changed largely from 0.5 to 92.6 μg m<sup>-3</sup>  
203 during the measurement period, and accounted for 3% to 58% of total WSI (=45.7 +  
204 30μg m<sup>-3</sup>). The highest hourly nitrate concentration occurred on December 23, 2015  
205 together with high concentrations of sulfate (65.5 μg m<sup>-3</sup>), and ammonium (56.8 μg m<sup>-3</sup>).  
206 Heavy haze episode like this occurred at our site frequently during winter. To  
207 understand the influence of wind on nitrate concentration, the wind rose plot is given  
208 in Fig. S2. The prevailing winds at the SORPES station were from northeast and east  
209 during the two-year observation period. Particulate nitrate tended to accumulate or  
210 formed under stagnant condition of low wind speed. The wind from west and east can  
211 lead to a higher nitrate concentration and also other aerosol components (Ding et al.,  
212 2013bc; Shen et al., 2018), which may be associated with air masses from biomass  
213 burning region and the city clusters of YRD.

214 To get an overall picture of nitrate distribution in developed region of costal China,  
215 we reviewed results from available nitrate measurements in three most polluted regions  
216 of North China Plain (NCP), YRD and Pearl River Delta (PRD), and summarized their  
217 concentration, sampling sites and measurement techniques in Fig.1. The corresponding  
218 data and references are listed in table S2. Measurements in summer and winter were  
219 separated due to the large seasonal difference of particulate nitrate. Despite that these  
220 measurements were from various measurement techniques, the results still can give us  
221 some insights about the differences in spatial and temporal scales. First, particulate  
222 nitrate generally showed the highest concentration in NCP and followed by YRD and  
223 PRD. This was in consistence with the spatial distribution of NO<sub>2</sub> – a major gas  
224 precursor of nitrate. Second, evident seasonal variations can be observed at all three



225 regions with much higher concentrations in winter. Third, an overall increase of  
226 particulate nitrate was implied in NCP and YRD in the past decade, especially that  
227 during summertime (shown in Fig.1a). Nevertheless, particulate nitrate in PRD  
228 revealed an overall decreasing trend. It should be noted that the dataset cited in Fig. 1  
229 were obtained from different sites with different techniques. Trends inferred from these  
230 datasets could suffer from considerable uncertainty. Compared to these previous studies,  
231 the nitrate concentration during summertime at SORPES station was lower than that in  
232 NCP, but higher than that in YRD and PRD cities. In terms of wintertime, nitrate  
233 concentration at SORPES station was slightly lower than that in NCP, comparable with  
234 that in YRD, and higher than that in PRD.

235 Fig. 2 illustrates the occurrence frequency of the loading of particle matter in  
236 different concentration ranges, and the changes of nitrate proportion along with the PM  
237 loading. Noting that the PM loading here was indicated by the mass of total WSI. The  
238 highest frequency of WSI concentrations occurred in a range of 20-40  $\mu\text{g m}^{-3}$ , and  
239 gradually decreased with the increasing of concentration. Heavy PM pollution with  
240 WSI concentrations higher than 100  $\mu\text{g m}^{-3}$  occurred during more than 5% of the time  
241 during this study. The contribution of nitrate to total WSI increased with the PM loading,  
242 ranging from ~25% with WSI concentration lower than 20  $\mu\text{g m}^{-3}$  to ~ 40% when WSI  
243 was higher than 140  $\mu\text{g m}^{-3}$ . These results suggested that nitrate was a major driver of  
244 haze episodes with high PM peaks in this region.

245 Fig. 3a shows the scatter plot of particulate nitrate and total WSI. The correlation  
246 coefficient was 0.92 and nitrate accounted for 32% of the total WSI. Air temperature  
247 greatly affected the contribution of nitrate to total WSI. Its proportion can be up to 58%  
248 at in the temperature range of 0 °C to 5 °C and only 3% at the temperature higher than  
249 30 °C, indicating an important role of thermodynamic equilibrium in nitrate  
250 concentration. We further investigated the neutralization extent of sulfate and nitrate by  
251 ammonium (Fig. 3b). Ammonium was overall enough to neutralize both sulfate and  
252 nitrate, suggesting that the particulate nitrate mostly existed as ammonium nitrate at

253 SORPES station. This is different from the ammonia poor regions, where the uptake of  
254  $\text{HNO}_3$  to dust/sea salt particles was found to be important to  $\text{PM}_{2.5}$  nitrate (Griffith et  
255 al., 2015). Seasonal difference can be observed for the molar ratio of ammonium to the  
256 sum of sulfate and nitrate. In spring and early summer, a fraction of the particulate  
257 nitrate is present as  $\text{Ca}(\text{NO}_3)_2$  and  $\text{KNO}_3$ , which could explain the data points below the  
258 regression line in Fig. 3b. In winter, considerable ammonium is existed as  $\text{NH}_4\text{Cl}$  (Hu  
259 et al., 2017), resulting in the data points above the regression line.

### 260 3.2. Characteristics of fine particular nitrate in different seasons.

#### 261 3.2.1. Seasonal pattern and its main causes

262 Fig. 4 shows the composite seasonal pattern of  $\text{NO}_x$ ,  $\text{PM}_{2.5}$  nitrate, sulfate and the  
263 molar-based ratio of nitrate to sulfate during the 2-year period at SORPES station. The  
264 seasonal variations of other related species and nitrate-to- $\text{PM}_{2.5}$ -ratio are shown in Table  
265 S3. Similar to the previous studies (Griffith et al., 2015), a typical seasonal variation  
266 was observed for particulate nitrate (and its ratio to sulfate, i.e.  $\text{NO}_3^-/\text{SO}_4^{2-}$ ), with a  
267 maximum value of  $23.7 \mu\text{g m}^{-3}$  (140%) in January, and a minimum of  $8.4 \mu\text{g m}^{-3}$  (66%)  
268 in August and September. Particulate sulfate exhibits a relatively less pronounced  
269 seasonal pattern with a small peak in June. The low value of particulate nitrate during  
270 summer can be generally explained by the higher temperature, higher and unstable  
271 boundary layer and relative clean air masses induced by summer monsoon (Ding et al.,  
272 2013c) despite of the increased photochemical formation. In opposite, the high values  
273 during winter were generally due to the lower temperature, lower and stable boundary  
274 layer and relative stronger continental outflow from the North China where  
275 anthropogenic emission was relatively high due to heating in winter (Ding et al., 2013c).  
276 Different chemical processes that affects nitrate concentrations between summer and  
277 winter will be discussed later.  $\text{NO}_x$ , the major precursor, tracked the changes of  
278 particulate nitrate, except for that during February and June. In addition, a secondary  
279 peak of particulate nitrate can be observed during June, which can be explained as the  
280 influence from agricultural burning in eastern China (Ding et al., 2013bc; Xie et al.,

281 2015; Shen et al., 2018). The concentrations of Potassium, a biomass burning tracer in  
282 this region (Ding et al., 2013b; Xie et al., 2015), clearly showed a consistent peak (Fig.  
283 S3) with both particulate nitrate and sulfate, as well as the discrepancy of NO<sub>x</sub> and  
284 nitrate concentrations (Ding et al., 2013b; Xie et al., 2015; Nie et al., 2015) shown in  
285 Fig. 4. While in February, the nitrate concentration didn't show concurrent decrease in  
286 NO<sub>x</sub> during the Chinese Spring Festival (Ding et al., 2013c). The observations might  
287 suggest that particulate nitrate is influenced by regional transport but not the local  
288 emissions in February.

### 289 3.2.2. Diurnal cycles during summer and winter

290 In Fig. 5, we show the averaged diurnal variations of particulate nitrate, nitrogen  
291 dioxide, nitric acid, equilibrium constant ( $K$ ), air temperature and RH during summer  
292 and winter during the two years. Nitric acid was calculated by ISORROPIA II. In  
293 summer (Fig. 5a), the fine particulate nitrate showed a typical diurnal cycle that the  
294 maximum concentration occurred at 7:00 with the average concentration of 16.5  $\mu\text{g m}^{-3}$   
295 and minimum value at 14:00 (7.2  $\mu\text{g m}^{-3}$ ). This summertime diurnal pattern of nitrate  
296 is very similar with the findings in Shandong (Wen et al., 2015) and New York (Sun et  
297 al., 2011). However, it is quite different from the findings in Hong Kong (Griffith et al.,  
298 2015), where nitrate concentration peaks in the daytime in summer. Ambient  
299 temperature and the development of boundary layer are the major drivers to the  
300 observed diurnal variation of particulate nitrate, and high temperature and high  
301 boundary layer during daytime lead to evaporation and dilution of the particulate nitrate  
302 (Zhang et al., 2015a; Ding et al., 2016). Nitric acid, which accounted for 20% of the  
303 total nitrate [ $\text{NO}_3^- + \text{HNO}_3$ ], revealed its high concentration (around 2 ppb) in the  
304 noontime (12:00-15:00). NO<sub>2</sub>, the precursor of nitrate, showed a peak concentration of  
305 18.2 ppb at 21:00, and remained at a high level during the whole night. The Equilibrium  
306 constant,  $K$ , was calculated to understand the influence of gas-to-particulate partitioning  
307 on the observed diurnal variation of particulate nitrate (Sun et al., 2011). As showed in  
308 Fig. 5a,  $K$  showed the same diurnal pattern as particulate nitrate, suggesting the

309 thermodynamic is the major factor influencing the diurnal variation of particle nitrate  
310 during summer.

311 In winter, the diurnal variation is small with a moderate peak at around 10:00 AM.  
312 Compared to that in summer, K showed similar diurnal variation, but not correlated to  
313 particulate nitrate, indicating factors other than the control of temperature. The  
314 observed peak at late morning was probably due to downward mixing from the residual  
315 layer where particulate nitrate was formed aloft during the night and brought to the  
316 surface after sunrise following the breakup of the boundary layer (Baasandorj et al.,  
317 2017; Young et al., 2016; Pusede et al., 2016). Direct vertical observations are needed  
318 to further investigate this issue.

319 To further investigate factors influencing the nitrate behaviors other than  
320 thermodynamics, ISORROPIA II was used to simulate the diurnal variation of nitrate.  
321 Hourly concentrations of all species (both gas and aerosol phase species) at 00:00 were  
322 used as the initial value of each specific day. Hourly data of temperature and relative  
323 humidity were used as the input data to constrain the model. The ISORROPIA II model  
324 was set as forward mode and metastable phase state. The calculated diurnal variations  
325 are shown in Fig. 6 together with the observed results.

326 The differences between the calculation and the observation could be attributed to:  
327 (1) the development of boundary layer, (2) the dry deposition of nitric acid, and (3)  
328 chemical processes, which has not been considered yet in the model. As shown in Fig.  
329 6a, the overall diurnal pattern of nitrate in summer is well captured by the model except  
330 for 3 periods. The differences after midnight are likely caused by the effect of boundary  
331 layer height and some chemical processes. Faster increase of model nitrate after 18:00  
332 was attributed to lack of dry deposition of nitric acid in the model. During noontime the  
333 observed nitrate concentration was expected to be lower than the calculated value  
334 because of the development of boundary layer and stronger dry deposition of nitric acid  
335 associated with stronger turbulence mixing, which were neglected in the model.  
336 However, in contrast, the observation was considerably higher than the calculated value.

337 It indicates a strong production of nitrate via photochemical processes in summer.  
338 Fig.6c shows that the difference between calculated and observed nitrate concentration  
339 was in good correlation with the product of NO<sub>2</sub> and solar radiation, a proxy for the  
340 production rate of nitric acid (Zhang et al., 2005; Young et al., 2016), further suggesting  
341 that photo-oxidation of NO<sub>2</sub> is an important source of nitrate during summer, even  
342 though the thermodynamic equilibrium is the dominate factor controlling the diurnal  
343 cycle. Wen et al. (2018) also demonstrated that photochemical production of nitric acid  
344 is a major contributor to daytime nitrate increase during summer in North China Plain.  
345 In winter, the influence of thermodynamics on diurnal cycle is small since the relatively  
346 low temperature throughout the day. The peaks in the morning may be caused by the  
347 mixing down of a residual layer enriched of nitrate as mentioned above, while the  
348 decline during the afternoon are supposed to be the result of dilution associated with  
349 the boundary layer development.

### 350 3.2.3 Influence of air masses transport

351 Meteorological processes play a key role in air masses long-range transport and local  
352 accumulation (Ding et al., 2013ac; Zhang et al., 2016; Ding et al, 2016b). In order to  
353 investigate the influence of air masses transport on nitrate concentrations, Lagrangian  
354 dispersion modeling was conducted for the sampling days with the top and bottom 25%  
355 nitrate concentration in summer and winter, respectively (Ding et al., 2013a). Fig. 7  
356 shows the retroplumes, i.e. footprint at an altitude of 100 m, of the selected days during  
357 summer and winter, respectively. In summer, high concentrations of nitrate tended to  
358 be associated with the air masses from west of Nanjing (mostly Anhui province) and  
359 Yangtze River Delta (Suzhou-Shanghai city clusters and North Zhejiang province) (Fig.  
360 7a). YRD is a high NO<sub>x</sub> emission region (Fig. 1), air masses from which could bring  
361 high concentration of NO<sub>x</sub> to enhance the nitrate concentration at SORPES station.  
362 Biomass burning is the possible cause of the high nitrate loading with air mass from the  
363 west of Nanjing (Fig. S4). In winter, regional transport from northern China played an  
364 important role in enhancing nitrate concentrations. As shown in Fig. 8c, a large part of

365 air masses for the highest 25% sampling days was from North China Plain, which has  
366 the strongest NO<sub>x</sub> emission in China (Fig. 1). It should be noted here that the longer  
367 lifetime of particulate nitrate during winter might be the main cause to promote the  
368 contribution of regional transport to the observed nitrate at SORPES. In contrast, the  
369 lowest 25% sampling days during winter tended to be accompanied with the air mass  
370 from Nanjing local and marine areas.

### 371 3.3. Contribution of N<sub>2</sub>O<sub>5</sub> hydrolysis to nitrate episodes

372 Similar to findings from previous studies (Zhang et al., 2015c), nitrate was found to  
373 increase significantly during this study and become the largest contributor of PM<sub>2.5</sub>  
374 during the haze episodes (Fig. 2). Generally, these pollution episodes mainly occurred  
375 in winter (Fig. 3a and Fig. 4), during which the photochemical production of nitric acid  
376 should be weak. N<sub>2</sub>O<sub>5</sub> hydrolysis was thus proposed to be a potential important  
377 formation pathway. Here we investigated the nitrate episodes in detail and discussed  
378 their relationship to the N<sub>2</sub>O<sub>5</sub> hydrolysis during the nights before.

379 In Fig. 8, we show a typical case of nitrate episodes from 30 November to 2  
380 December, 2015. Fast nitrate formation was observed, which was likely caused by  
381 hydrolysis of N<sub>2</sub>O<sub>5</sub>. Nitrate increased significantly from 20.3 μg m<sup>-3</sup> at 18:00 of 30  
382 November to 63 μg m<sup>-3</sup> at 6:00 of 1 December, 2015. The ratio of nitrate to PM<sub>2.5</sub> also  
383 exhibited a large increase from 25% at 18:00 to 38% at 06:00. In contrast, other PM<sub>2.5</sub>  
384 components, e.g. sulfate and black carbon, showed only slight increases. High  
385 concentration of NO<sub>2</sub>, considerable level of O<sub>3</sub> and extremely low concentration of NO  
386 provided a favorable condition towards forming NO<sub>3</sub> and N<sub>2</sub>O<sub>5</sub> (Brown et al., 2003).  
387 The meteorological conditions during these 12 hours were stable with low wind speed  
388 and high relative humidity, which, combined with the relatively high concentration of  
389 PM, would promote the hydrolysis of N<sub>2</sub>O<sub>5</sub> (Riemer, 2003).

390 N<sub>2</sub>O<sub>5</sub> concentrations were calculated by using the steady-state approximation  
391 (Osthoff et al., 2006; Wang et al., 2014; Wen et al., 2015), and the result was shown in  
392 Fig. 8. The calculated N<sub>2</sub>O<sub>5</sub> exhibited a much higher concentration during the night of

393 30 November compared to the days before and after. Particulate nitrate formed from  
394  $\text{N}_2\text{O}_5$  hydrolysis was then computed during the 12-hour period. Nitrate concentration  
395 at 18:00 of 30 November, 2015 ( $20 \mu\text{g m}^{-3}$ ) was selected as the initial value, and  $31 \mu\text{g}$   
396  $\text{m}^{-3}$  of particulate nitrate was produced in the following 12 hours, suggesting that  
397 approximate 80% of increased particulate nitrate can be attributed to the hydrolysis of  
398  $\text{N}_2\text{O}_5$  in this case.

399 To further understand the contribution of  $\text{N}_2\text{O}_5$  hydrolysis, sampling days with  
400 daily-averaged nitrate concentration exceeding the mean plus twice the standard  
401 deviation were selected as the nitrate episode days. In total, 18 episode-days were  
402 selected during the 2-year measurement, with 16 days in winter and the other 2 days in  
403 biomass burning season. In Fig. 9, we presented the averaged diurnal pattern of  
404 particulate nitrate and its related parameters on the 18 selected episode and pre-episode  
405 days. For the episode days, particulate nitrate revealed a similar diurnal pattern as that  
406 of the whole winter (Fig. 5). Nitrate maintained a high concentration during the whole  
407 day with a small peak around 10:00 in the morning. However, for the pre-episode days,  
408 a clear build-up of nitrate can be observed, especially during the nighttime from 17:00  
409 of the pre-episode days to 1:00 of the episode days (as marked in Fig. 9). The average  
410 increment of ammonium nitrate exceeded  $24 \mu\text{g m}^{-3}$  during this 9-hour period of the  
411 pre-episode nighttime. The total WIS also increased during this period, which was  
412 mostly attributed to ammonium nitrate (almost 90%) and resulted in an evident increase  
413 of the ratio of nitrate to total WSI. Compared to nitrate, black carbon, a tracer of primary  
414 emissions, showed little change during the pre-episode day. The retroplume showed in  
415 Fig. S5 suggested that the air masses arrived at the SORPES station on the pre-episode  
416 and episode days were almost the same. These results suggest that secondary formation  
417 other than accumulation was the major contributor to the observed increase of  
418 particulate nitrate.

419 Since the observed nitrate formation mostly occurred during the nighttime of pre-  
420 episode days when the photochemical production of nitric acid would be largely

421 suppressed,  $\text{N}_2\text{O}_5$  hydrolysis is thus believed to be the major contributor. As showed in  
422 Fig. 9, compared to those during episode days,  $\text{NO}_2$  concentration was comparable with  
423 the average concentrations of 28 ppb, the but  $\text{O}_3$  concentration was higher during pre-  
424 episode days. This resulted in a higher production rate of  $\text{N}_2\text{O}_5$  proxy ( $\text{PNO}_3$ ) in pre-  
425 episode days, and favored formation of nitrate from the hydrolysis of  $\text{N}_2\text{O}_5$ . We further  
426 calculated the contribution of  $\text{N}_2\text{O}_5$  hydrolysis to nitrate formation during the periods  
427 from 17:00 to 23:00 of each pre-episode day (excluding 2 windy days). A good  
428 correlation ( $R=0.8$ ) was observed between the calculated nitrate and observed nitrate  
429 (Fig. 10), with the slope of 0.77, indicating the observed nitrate formation during nitrate  
430 episodes were significantly attributed to the hydrolysis of  $\text{N}_2\text{O}_5$ . It should be noted that  
431 this calculation suffered from considerable uncertainties due to the variability of actual  
432 VOCs concentrations and  $\text{N}_2\text{O}_5$  uptake coefficient. The detailed uncertainty calculation  
433 is discussed in the supplement (Fig. S1).

#### 434 **4. Summary and Conclusion**

435 Online measurements of fine particulate nitrate along with trace gases and  $\text{PM}_{2.5}$  mass  
436 concentrations were conducted for two years from March 2014 to February 2016 using  
437 a MARGA at SORPES station, a rural receptor site in the Yangtze River Delta, eastern  
438 China. Hourly nitrate concentration varied from  $0.5 \mu\text{g m}^{-3}$  to  $92.6 \mu\text{g m}^{-3}$ , with an  
439 averaged value of  $15.8 \mu\text{g m}^{-3}$ , which was generally higher than the measurement at the  
440 sites in YRD and PRD, but lower than that at the sites in North China Plain. The  
441 contribution of nitrate to total WSI increased from 25% with WSI concentration lower  
442 than  $20 \mu\text{g m}^{-3}$ , to 40% when WSI was higher than  $140 \mu\text{g m}^{-3}$ , suggesting a major  
443 driver of nitrate to the aerosol pollution in YRD.  $\text{NH}_3$  is enough to neutralize the acidic  
444 compounds of aerosol, and ammonium nitrate was thus the predominate form of the  
445 observed particulate nitrate. A clear seasonal variation of nitrate was observed with  
446 peak value in January and December and lowest value in August and September.  
447 Biomass burning plumes contributed to the nitrate concentration evidently and resulted  
448 in a secondary peak during June. In summer, thermodynamic equilibrium was the major



449 factor influencing the diurnal variation of nitrate, and resulted in a much lower  
450 concentration at noontime. Nevertheless, the observed nitrate at noontime was  
451 considerably higher than the value predicted by the ISORROPIA II model, indicating a  
452 strong production of nitrate by the photo-oxidation of  $\text{NO}_2$ . Air masses from YRD and  
453 the biomass burning region were associated with the high nitrate concentrations during  
454 summer. In winter, the diurnal variation of nitrate was weak. Regional transport from  
455 the North China Plain contributed largely to the observed high nitrate concentrations.

456 Nitrate episodes, defined as daily-averaged concentration exceeding the mean  
457 value plus twice the standard deviation, were further investigated to understand the  
458 chemical processes towards forming particulate nitrate and their contribution to the  
459 pollution episodes. A clear build-up of nitrate can be observed during the pre-episode  
460 night, which dominated the increase of total WSI.  $\text{N}_2\text{O}_5$  hydrolysis was demonstrated  
461 to contribute 80% of the observed nitrate formation, suggesting its critical role in an  
462 aerosol pollution episode. In view of the significant emission of  $\text{NO}$ , which is the main  
463 sink of  $\text{N}_2\text{O}_5$  during night, stronger production of  $\text{N}_2\text{O}_5$  is expected at the upper  
464 boundary layer, e.g. residual layer, and this residual layer  $\text{N}_2\text{O}_5$  will contribute to the  
465 nitrate formation in the entire boundary layer. In summary, our study provides evidence  
466 that particulate nitrate especially that formed from  $\text{N}_2\text{O}_5$  hydrolysis is a crucial  
467 contributor to the aerosol pollution episodes in eastern China.

468 *Data availability.* The GDAS data used in the HYSPLIT calculation can be acquired  
469 from <ftp://arlftp.arlhq.noaa.gov/pub/archives/gdas1>. Measurement data at SORPES,  
470 including aerosol data and relevant trace gases as well as meteorological data, are  
471 available upon request from the corresponding author before the SORPES database is  
472 opened publicly.

473 *Acknowledgements.* The research was supported by National Key Research &  
474 Development Program of China (2016YFC0200500) and National Science Foundation

475 of China (41725020, 91544231, 91644218, 41422504, 91744311, 41675145).

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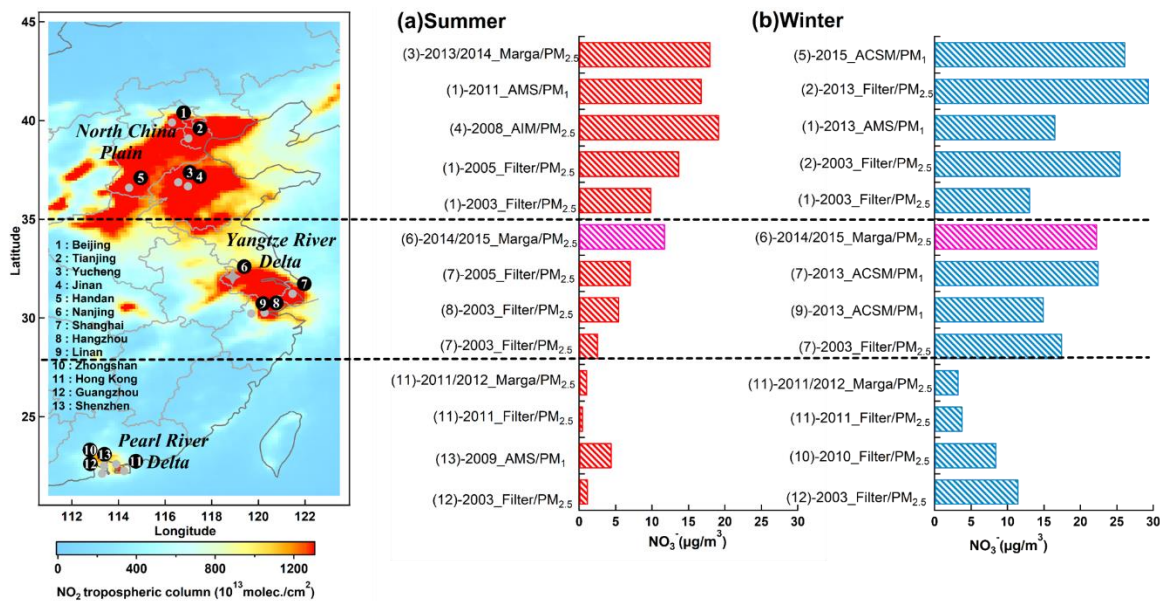
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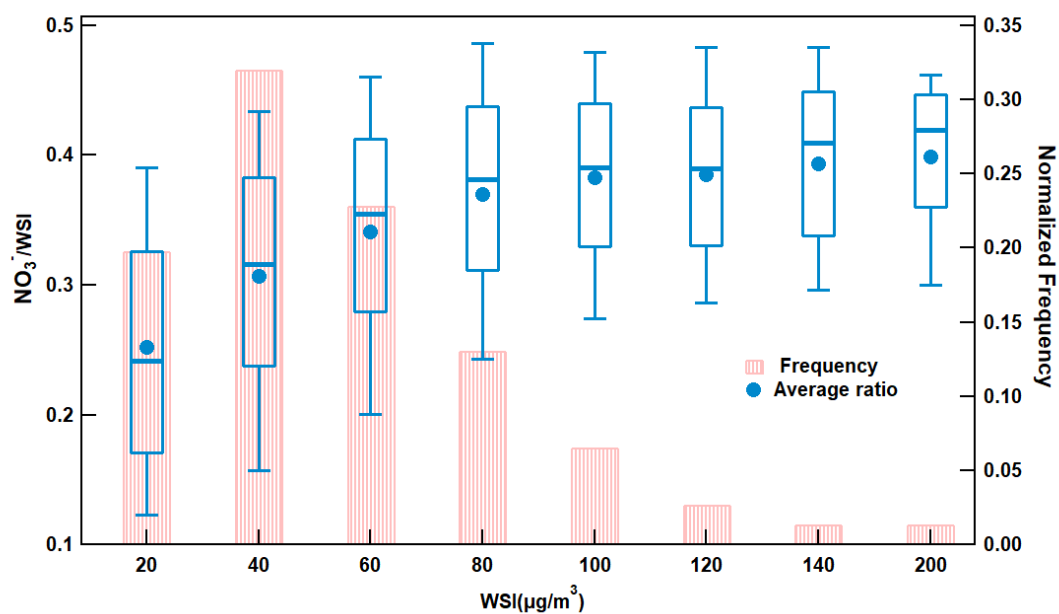
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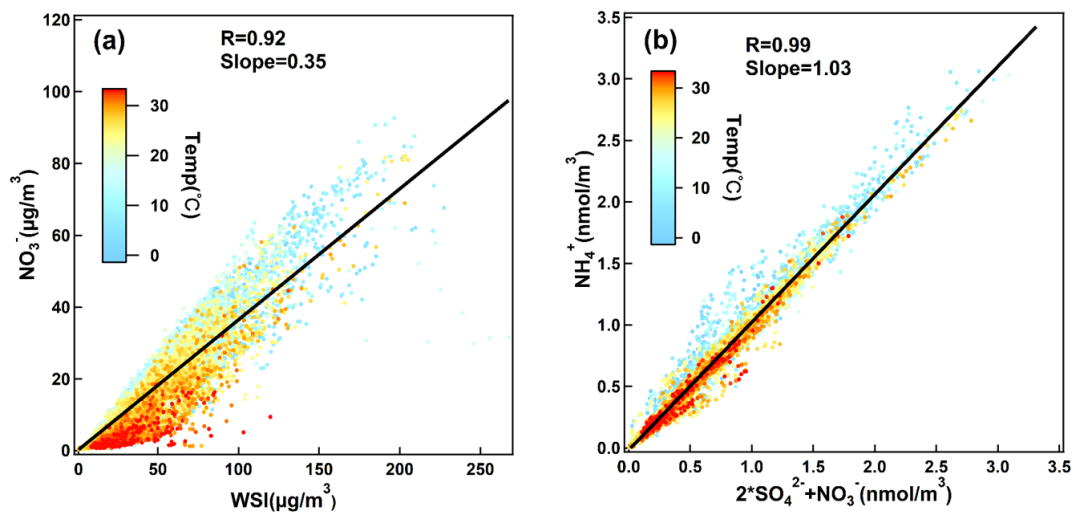




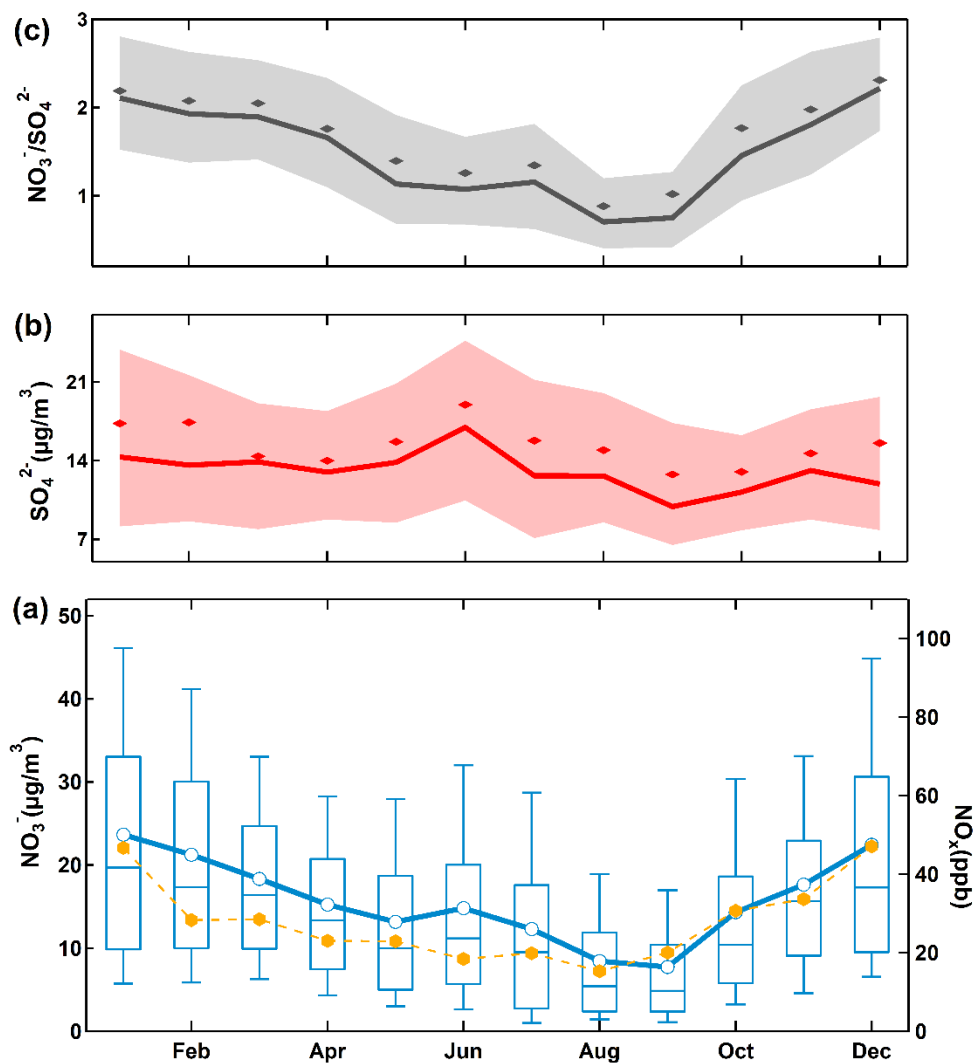
**Figure 1** Average mass concentrations of particulate nitrate at different sampling sites in **(a)** summer and **(b)** winter. The left panel shows the map color-coded by 2-years (2014-2015) averaged tropospheric NO<sub>2</sub> from OMI satellite (<http://www.temis.nl/airpollution/no2.html>). The pink bars are for this study.



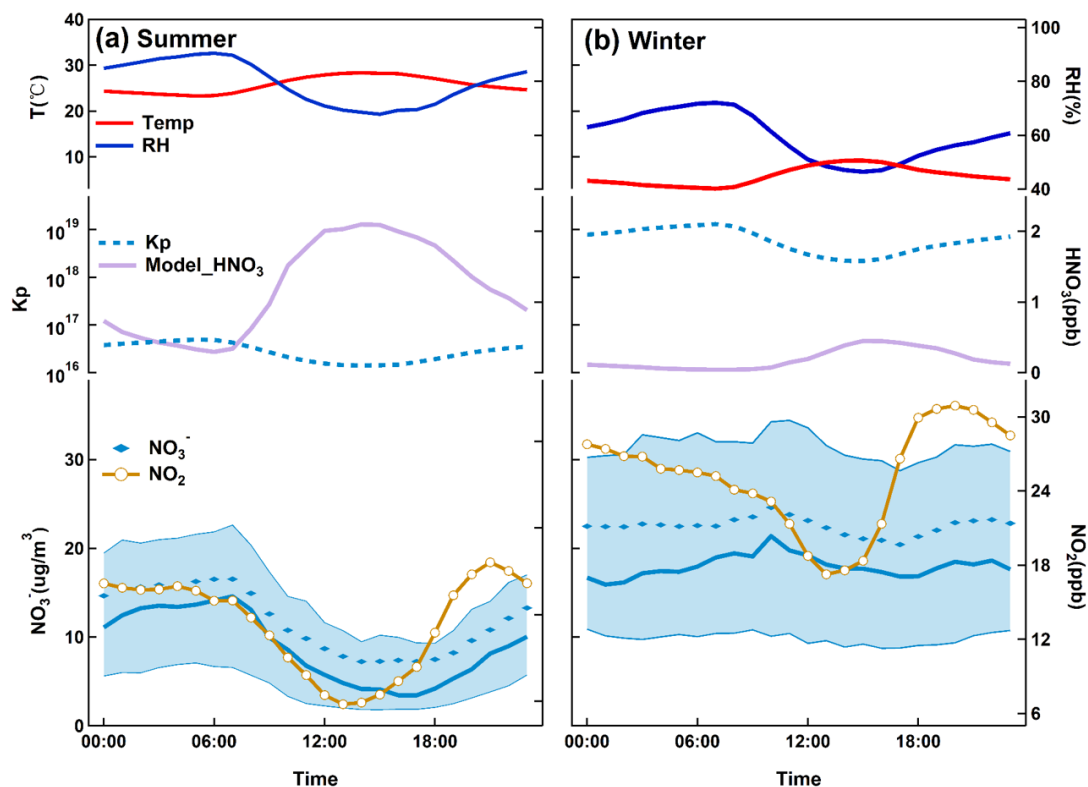
**Figure 2** Average proportion of nitrate and normalized frequency of occurrence at different mass concentration bins of water soluble ions at SORPES. For the ratio, box boundaries represent the interquartile range, bars represent 5%-95% percentile range, and horizontal lines represent the median value.



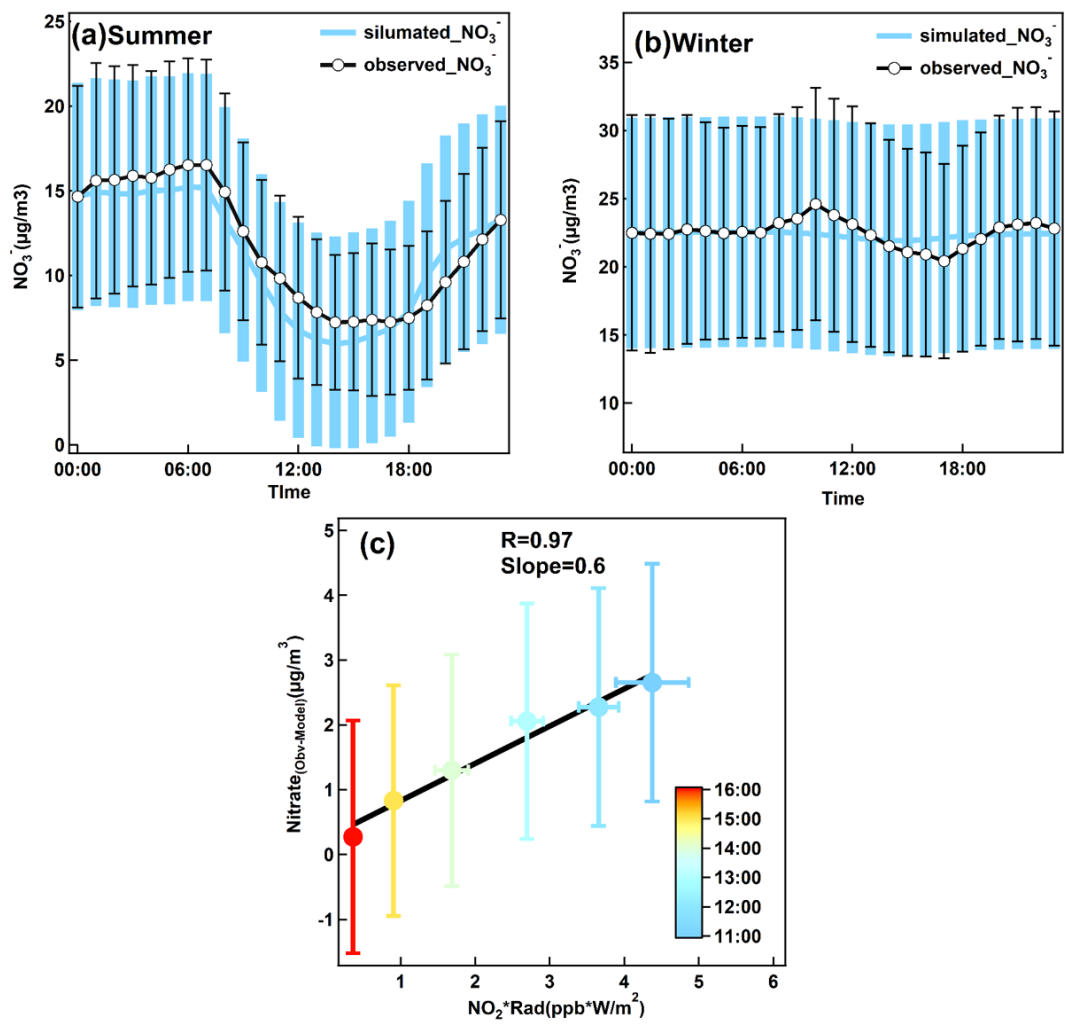
**Figure 3** Scatter plots of (a) nitrate vs. total WSI color coded by air temperature, (b) molar concentrations of ammonium with nitrate molar concentrations plus two times of sulfate molar concentrations.



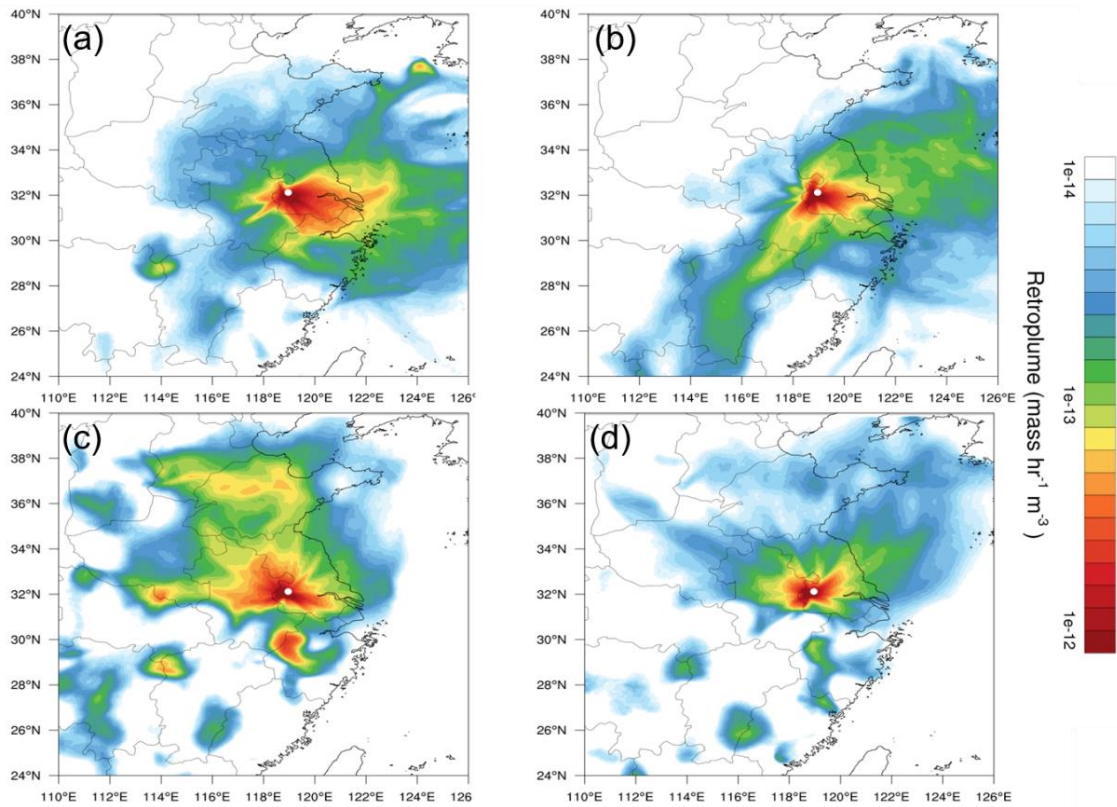
**Figure 4** Monthly averaged nitrate (blue), sulfate (red),  $\text{NO}_x$  (orange) mass concentrations and nitrate to sulfate molar-based ratio (grey) measured at SORPES station during March 2014 to February 2016. For nitrate to sulfate ratio (a) and sulfate (b), bold solid lines are the median values, shade areas represent percentiles of 75% and 25%, and diamonds represent the mean values. For nitrate (c), box boundaries represent the interquartile range, bars represent 5%-95% percentile range, horizontal lines represent the median value, and crosses represent mean values.



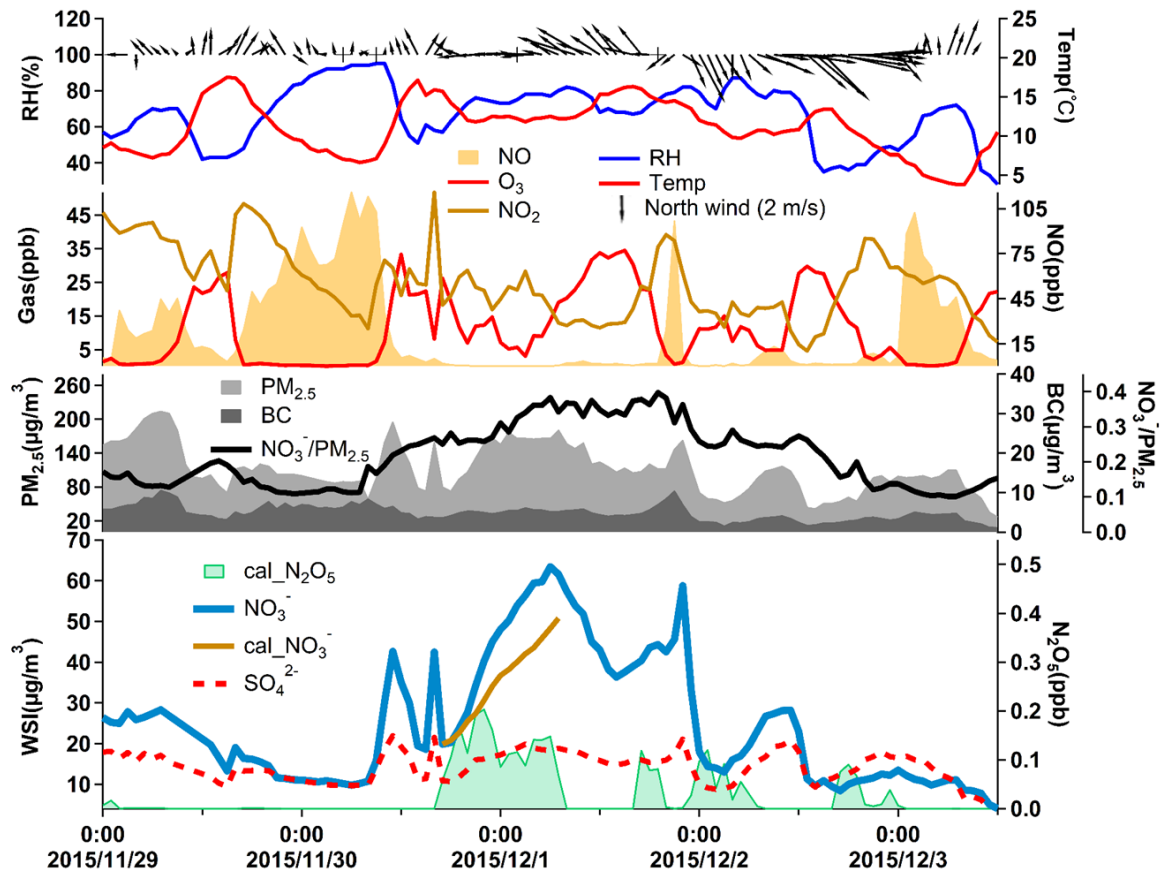
**Figure 5** Diurnal variation of particulate nitrate in **(a)** summer and **(b)** winter. For nitrate, bold solid lines are the median values, shaded areas represent percentiles of 75% and 25% and diamonds represent mean values. Diurnal averages of  $\text{NO}_2$  and modeled nitric acid concentrations are also provided with temperature and RH.



**Figure 6** Modeled nitrate diurnal variations in **(a)** summer and **(b)** winter, together with the observed nitrate concentrations. Error bars provided are the standard deviation of the mean at each hourly interval. **(c)** Scatter plot of the difference between model and observed nitrate average mass concentrations with the product of NO<sub>2</sub> and radiation color coded by the hour of day for the samples.

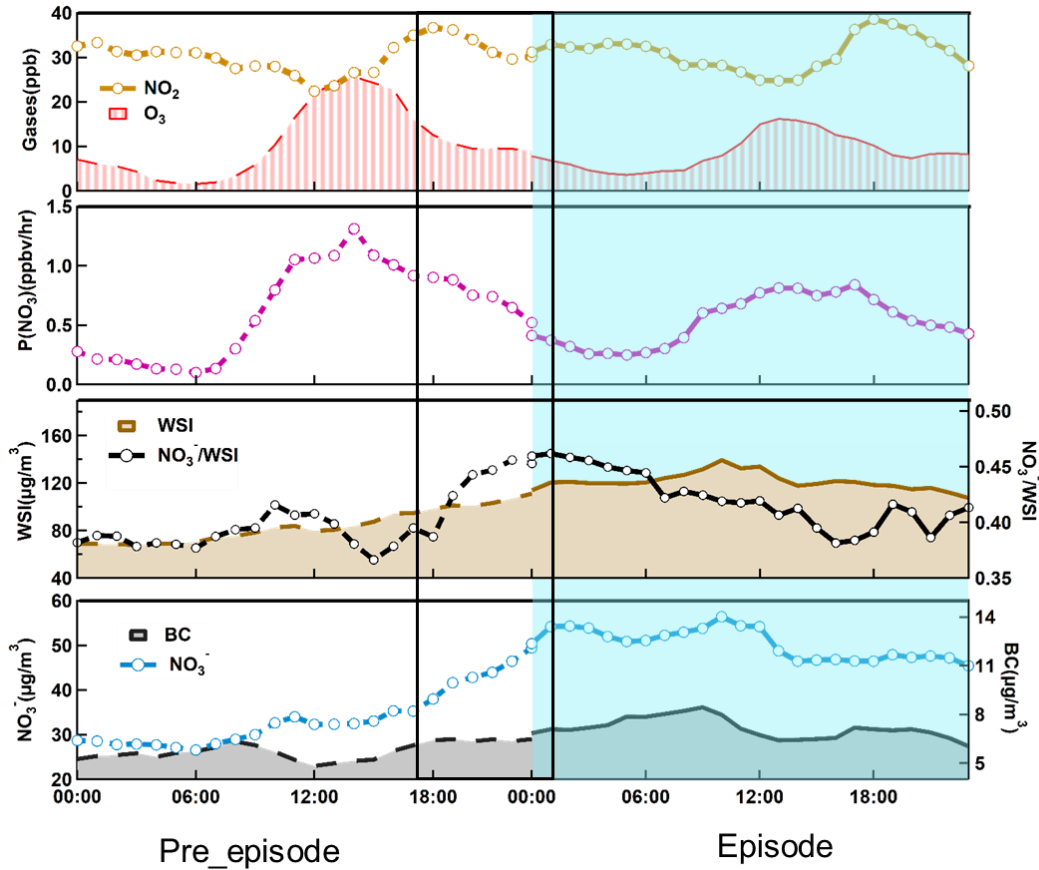


**Figure 7** The averaged retroplumes (i.e., 100 m footprint) of the selected events: **(a)** Top 25% nitrate concentrations in summer, **(b)** Bottom 25% nitrate concentrations in summer, **(c)** Top 25% nitrate concentrations in winter, and **(d)** Bottom 25% nitrate concentrations in winter.

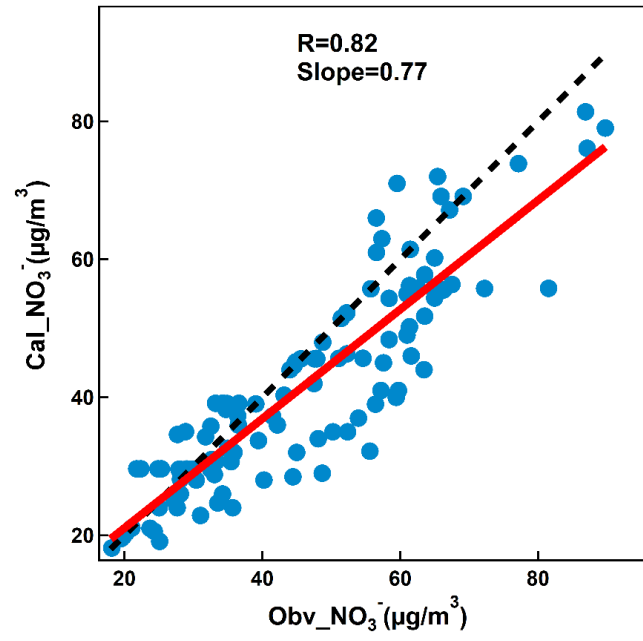


**Figure 8** Time series of meteorological data and the concentrations of trace gases related to nitrate formation during 29 November to 3 December, 2015. Cal\_NO<sub>3</sub><sup>-</sup> represents the nitrate concentrations calculated from the hydrolysis of N<sub>2</sub>O<sub>5</sub>.





**Figure 9** Diurnal variations of particulate nitrate, black carbon, the total water soluble ions, nitrate to WSI ratio,  $P(\text{NO}_3)$ ,  $\text{NO}_2$ , and  $\text{O}_3$  averaged for nitrate episode days with exceedances of one mean plus two standard deviations. The left side shows the pre-episode days and the right side shows the episode days during the winter of entire two years period. The solid line box corresponds to the rapid growth of nitrate at night. It should be noted that  $P(\text{NO}_3)$  is calculated by the product of  $\text{NO}_2$  and  $\text{O}_3$  multiplied by the rate constant  $k_1$  of  $\text{NO}_2 + \text{O}_3$  reaction.



**Figure 10** Scatter plot of calculated nitrate concentrations and observed nitrate concentrations from 17:00 to 23:00 of each episode.

Table 1 major gas phase and heterogeneous reactions involved NO<sub>3</sub> and N<sub>2</sub>O<sub>5</sub>

Reaction	Rate constant
$\text{NO}_2 + \text{O}_3 \rightarrow \text{NO}_3 + \text{O}_2$	$k_1 = 1.28 \times 10^{13} \times \text{EXP}(-2470/T)$
$\text{NO}_3 + \text{NO}_2 \leftrightarrow \text{N}_2\text{O}_5$	$k_{eq} = 1.73 \times 10^{-13} \times \text{EXP}(1550/T)$
$\text{NO}_3 + \text{NO} \leftrightarrow \text{NO}_2 + \text{NO}_2$	$k_3 = 1.8 \times 10^{-11} \times \text{EXP}(110/T)$
$\text{NO}_3 \rightarrow \text{NO} + \text{O}_2$	$j_4$
$\text{NO}_3 \rightarrow \text{NO} + \text{O}_2$	$j_5$
$\text{NO}_3 \xrightarrow{\text{voc}} \text{products}$	$k_6 = \sum (k_{\text{voc}i} \cdot [\text{voc}]_i)$
$\text{NO}_3 \xrightarrow{\text{Heterogeneous}} \text{products}$	$k_7 = 0.25 \cdot i \cdot C_{\text{NO}_3} \cdot \gamma_{\text{NO}_3} \cdot S_{\text{aerosol}}$
$\text{N}_2\text{O}_5 \xrightarrow{\text{Heterogeneous}} \text{products}$	$k_8 = 0.25 \cdot C_{\text{N}_2\text{O}_5} \cdot \gamma_{\text{N}_2\text{O}_5} \cdot S_{\text{aerosol}}$