

Supplementary Material

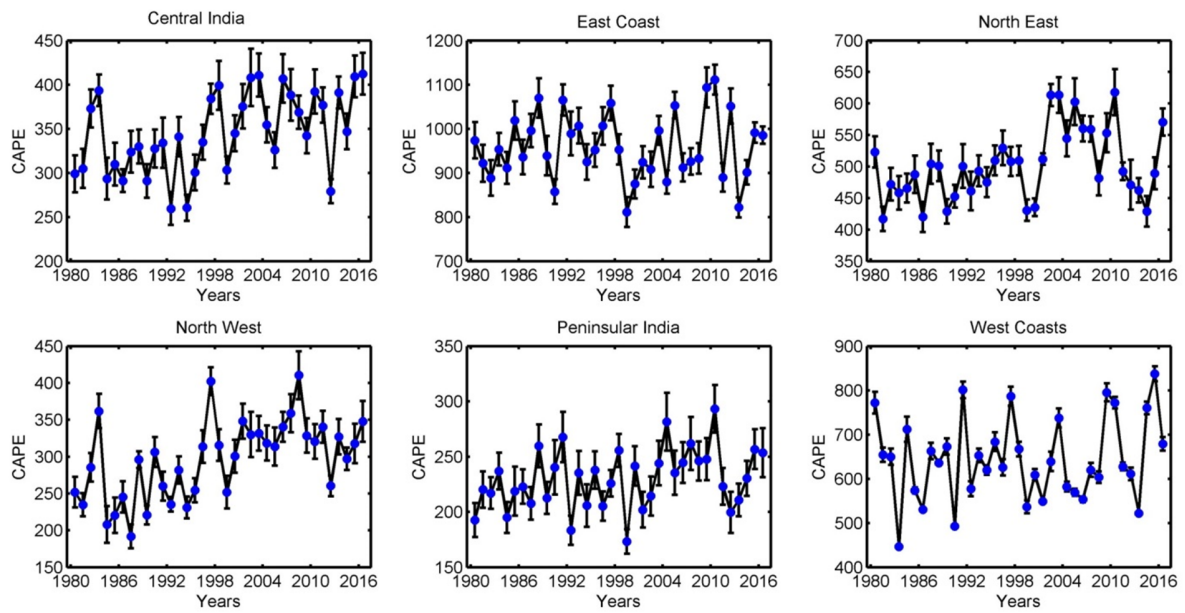


Figure S1: Climatic trends of CAPE for all six Indian Regions using ERA Interim monthly average datasets

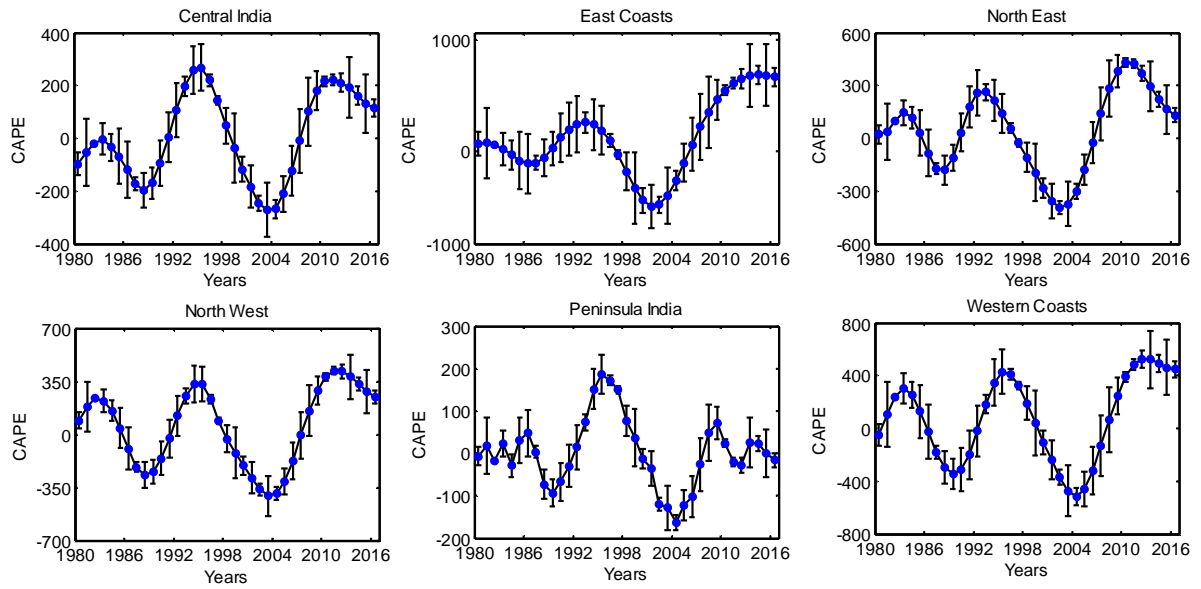


Figure S2: MCO periodicities of CAPE for all Indian regions.

1 **Text S1: Breif Description of the parcel and instability parameters used**

- 2 (a) Lifted Condensation level (LCL): It is the height at which an air parcel would attain condensation if it
3 is raised dry adiabatically from the surface. This is because at that height, the air parcel mixing ratio
4 becomes equal to saturation mixing ratio. Lower values of it indicate presence of tall clouds reaching the
5 surface during convective events.
- 6 (b) Level of Free condensation (LFC): It is defined as that level of atmosphere above the LCL where the
7 parcel of air lifted moist adiabatically would for the first time be warmer than its surrounding. Lower
8 heights of LFC indicate more probability of moist air parcels to be lifted upwards leading to more
9 convective strength in the system.
- 10 (c) Equilibrium Level (EL): It is the level of atmosphere where a parcel would attain same temperature as
11 of surroundings and would cease its upward motion. Higher the EL, more the energy contained within the
12 convective system.
- 13 (d) Lifted Index (LI): It is the difference between the temperature of the environment with that of an air
14 parcel lifted adiabatically to 500 mbar pressure. Negative or near negative value of LI reflects
15 condensation of saturated vapour parcel to liquid indicating instability and severe weather conditions. LI
16 generally indicate the intensity of the convective activities like CAPE.
- 17 (e) Vertical Totals Index (VTI): It represents the instability component of TTI. It is calculated as the
18 temperature gradient between 850 and 500 hPa pressure levels. This parameter is a sensitive measure of
19 atmospheric instability.
- 20 (f) Convective Available Potential Energy (CAPE): CAPE indicates the buoyant energy available to
21 accelerate an air parcel vertically and is calculated using the summation of positive buoyant energy from
22 LFC to EL. Higher CAPE provides more energy for convective growth; hence CAPE should be high in
23 convective conditions and less in normal conditions. According to some standard research attempts, the
24 values of CAPE > 1500 J kg⁻¹ are essential to have super cell formation.
- 25 (g) Mixed Layer CAPE (MLC): This parameter is similar to CAPE, but this one is mostly used to assess
26 the atmospheric instability during well mixed conditions in a better way than surface based CAPE. The
27 main difference between CAPE and MLCAPE is that, this parameter constitutes the total positive energy
28 attained by the parcel when lifted from the LFC to the lowest 100 mb of the troposphere. However, it
29 may be noted that in this paper, the effect of UTLS on CAPE and thunderstorm severity has been shown
30 in a focussed way. Also in clear weather conditions, EL mostly resides along ~300 hPa over all Indian
31 regions. Hence to separate the effect of UTLS near 100 hPa during intense convection, here the
32 summation is taken only upto 300 hPa levels.
- 33 (h) Convective Inhibition Energy (CINE): This is the summation of negative buoyant energy from surface
34 level to LFC. Being the opposite of CAPE, higher values of CINE produce strong hindrance to
35 convective genesis.
- 36 (i) Precipitable Water Vapor (PWV): This is the total column water vapour present per unit area. High
37 values indicate higher moisture favouring convective processes.
- 38 (j) Precipitable Water Vapour at Lower troposphere (PWL): Lower level moisture contents particularly
39 upto the boundary layer height (maximum 700 hPa) are also useful for controlling all rainfall occurrences

40 and hence they are taken additionally. Again, if both PWL and PWV are taken, then the contribution of
41 the middle and upper troposphere can be clearly understood for all Indian regions.

42 (k) Horizontal Wind Shear (WSH): This is a common parameter which is calculated as the gradient of
43 horizontal winds between the surface and the 6 km height level. Large values of this prevent convective
44 updrafts and hence it reduces thunderstorm severity.

45 (l) Temperature at 100 hPa Pressure levels (T100): This represents the atmospheric temperatures present
46 at the 100 hPa pressure level and it is recorded from the radiosonde profiles. This parameter is taken to
47 investigate whether at all there are any reliable relationships between UTLS cooling and CAPE as already
48 hinted earlier in previous research attempts over India.

49 (m) Ordinary Thunderstorm Frequency (TSO): This parameter is calculated for each station as the number
50 of radiosonde launches per year at 00Z where the collocated surface wind speed values as obtained from
51 base radionde data (obtained by Wyoming Portal is between 31-62 km/hr as already cited out by IMD
52 reports and Saha et al. 2018).

53 (n) Severe Thunderstorm Frequency (TSS): This is similar to TSO, only here the number of radiosonde
54 launches at 00Z are counted for which the surface wind speed is above 62 km/hr as again cited out by
55 IMD reports and Saha et al. 2018.

56 (o) Weak Rainfall Frequency (WRF): This parameter is calculated for each station as the number of
57 radiosonde launches or days per year at 00Z when the obtained one day rainfall accumulation is very low
58 (below 7.5 mm) as specified by IMD reports.

59 (p) Severe Rainfall Frequency (SRF): This parameter is similar to WRF; the only difference is that this
60 parameter is calculated for each station as the number of radiosonde launches or days per year at 00Z when
61 the obtained one day rainfall accumulation is very high (above 124.5 mm) as again cited by the IMD
62 reports.

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