

1 **Characteristics and mixing state of amine-containing**
2 **particles at a rural site in the Pearl River Delta, China.**

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43 **Abstract.** Particulate amines play an important role for the particle acidity and
44 hygroscopicity and also contribute to secondary organic aerosol mass. We investigated
45 the sources and mixing states of particulate amines using a single-particle aerosol
46 mass spectrometer (SPAMS) during summer and winter 2014 at a rural site in the
47 Pearl River Delta, China. Amine-containing particles accounted for 11.1 % and 9.4 %
48 of the total detected individual particles in summer and winter, respectively. Although
49 the increase of amine-containing particle count mostly occurred at night, no obvious
50 correlations between amine-containing particles and ambient relative humidity (RH)
51 were found during the sampling period. Among the three markers we considered, the
52 most abundant amine marker was $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$, which was detected in 90% and 86%
53 of amine-containing particles in summer and winter, followed by amine marker ions
54 of $^{59}(\text{CH}_3)_3\text{N}^+$, and $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$ which were detected in less than 10% of
55 amine-containing particles during sampling period. The amine-containing particles
56 were characterized by high fractions of carbonaceous marker ions, carbon-nitrogen
57 fragments, sulfate and nitrate in both summer and winter. More than 90% of
58 amine-containing particles were found to be internally mixed with sulfate throughout
59 the sampling period, while the percentage of amine particles containing nitrate
60 increased from 43% in summer to 69% in winter. Robust correlations between the
61 peak intensities of amines and the sum of nitrate and sulfate were observed,
62 suggesting the possible formation of aminium sulfate and nitrate salts. Interestingly,
63 only 8% of amine particles contained ammonium in summer, while the percentage
64 increased dramatically to 54 % in winter, indicating a relatively ammonium-poor state
65 in summer and an ammonium-rich state in winter. The total ammonium-containing
66 particles were investigated and showed a much lower abundance in ambient particles
67 in summer (3.6%) than that in winter (32.6%), which suggests the ammonium-poor
68 state of amine-containing particles in summer may be related to the lower abundance
69 of ammonia/ammonium in gas and particle phase. In addition, higher abundance of
70 amines in ammonium-containing particles than that of ammonium in
71 amine-containing particles suggests a possible contribution of ammonium–amine
72 exchange reactions to the low abundance of ammonium in amine-containing particles

73 at high ambient RH (72 ± 13 %) in summer. The particle acidity of amine-containing
74 particles is estimated via the relative acidity ratio (R_a), which is defined as the ratio of
75 the sum of the sulfate and nitrate peak areas divided by the ammonium peak area. The
76 R_a was 326 ± 326 in summer and 31 ± 13 in winter, indicating that the
77 amine-containing particles were more acidic in summer than in winter. However, after
78 including amines along with the ammonium in the acidity calculation, the new R_a '
79 values showed no seasonal change in summer (11 ± 4) and winter (10 ± 2), which
80 suggests that amines could be a buffer for the particle acidity of ammonium-poor
81 particles.

82 **Keywords:** Amine; Single particles; Mixing state; Aminium salts; Particle acidity;
83 SPAMS.

84 **1 Introduction**

85 Amines, a group of nitrogen-containing organic compounds, are ubiquitous in
86 the atmospheric gas and particle phases (Ge et al., 2011b). A variety of low molecular
87 weight (LMW) aliphatic amines have been detected in emissions from anthropogenic
88 and natural sources, including animal husbandry, biomass burning, industrial
89 emissions, vehicle exhaust, and marine sources (Rogge et al., 1994; Rappert and
90 Muller, 2005; Calderon et al., 2007; Ngwabie et al., 2007; Ge et al., 2011b). LMW
91 aliphatic amines have gas-phase concentrations two orders of magnitude lower than
92 that of ammonia (NH_3) (Sorooshian et al., 2008), but are more alkaline than NH_3 (Ge
93 et al., 2011a). Due to their strong basicity and water solubility, LMW amines can
94 undergo acid-base reactions with sulfuric and nitric acid to form aminium salts
95 (Angelino et al., 2001; Sorooshian et al., 2007; Pratt et al., 2009), which has been
96 found to enhance new particle formation beyond the amounts produced from reactions
97 between acids and NH_3 alone (Kurten et al., 2008; Berndt et al., 2010; Place et al.,
98 2010; Wang et al., 2010). In addition, once partitioned into the particle phase, these
99 LMW aliphatic amines can enhance aerosol particle hygroscopicity (Chu et al.,
100 2015; Sauerwein et al., 2015). Furthermore, amines can be oxidized by OH radicals,
101 NO_3 radicals, and O_3 in the atmosphere to form semi-volatile and non-volatile

102 compounds, some of which are highly toxic (Lee and Wexler, 2013), and which
103 contribute to secondary organic aerosol (SOA) mass (Murphy et al., 2007;Malloy et
104 al., 2009).

105 The mass concentration and temporal distribution of LMW aliphatic amines in
106 aerosols have been studied extensively in a variety of environments, and LMW
107 aliphatic amines account for 2–12 % of organic mass (Day et al., 2009;Gilardoni et al.,
108 2009;Liu et al., 2009;Russell et al., 2009;Williams et al., 2010). In recent years,
109 real-time single particle mass spectrometry has been used to measure the size and
110 chemical composition of individual amine-containing particles with high time
111 resolution. [The mixing state and single-particle characteristics of amines have been
112 investigated in laboratory and field environments \(Angelino et al., 2001;Moffet et al.,
113 2008;Silva et al., 2008;Pratt et al., 2009;Huang et al., 2012;Zhang et al., 2012\).](#) Pratt
114 et al. (2009) studied seasonal differences in aminium and ammonium salts on a
115 single-particle basis using an aerosol time-of-flight mass spectrometer (ATOFMS)
116 coupled with a thermodenuder and reported that the gas-to-particle partitioning of
117 amines is dependent on particle acidity. Healy et al. (2015) investigated the temporal
118 distributions of alkylamines at five European sites, and found that alkylamines were
119 internally mixed with both sulfate and nitrate, which suggests that the formation of
120 aminium salts was important at all sites. [Zauscher et al. \(2013\) detected strong signals
121 of amine marker \(\$^{86}\(\text{C}_2\text{H}_5\)_2\text{NCH}_2^+\$ \) in biomass burning aerosols associated with the
122 increase of ambient relative humidity, indicating the direct emission of amines from
123 biomass burning and the important influence of high RH \(>90%\) on the partitioning
124 process of amines.](#) Huang et al. (2012) determined the mixing state of
125 amine-containing particles in Shanghai and found higher number concentrations of
126 amine-containing particles in winter than in summer, which they attributed to
127 effective acid-base reactions between sulfuric acid and amines under low-temperature,
128 high-RH conditions. Zhang et al. (2012) measured trimethylamine-containing
129 particles in Guangzhou and found preferential trimethylamine gas-to-particle
130 partitioning during fog events. These field observations emphasize the important role
131 of acid-base reactions in the partitioning of amines from the gas phase to the particle

132 phase. Recent laboratory studies have revealed that the exchange between amine
133 gases and particulate NH_3 and/or ammonium also contributes substantially to amine
134 content and results in a depletion of NH_3 and/or ammonium in the particle phase
135 (Lloyd et al., 2009;Bzdek et al., 2010;Qiu et al., 2011;Liu et al., 2012;Chan and Chan,
136 2013;Chu and Chan, 2016, 2017;Sauerwein and Chan, 2017); however, the
137 significance of such exchange reactions in the ambient environment has not been fully
138 explored. Therefore, the influence of ammonia and particle acidity on the distribution
139 of amines in the particle phase should be studied comprehensively through field
140 measurements.

141 The aim of this study was to investigate the mixing state of a series of LMW
142 aliphatic amines with sulfate, nitrate, and ammonium in individual particles using a
143 single-particle aerosol mass spectrometer (SPAMS) at a rural site in the Pearl River
144 Delta, China. In order to explore amine origins and gas-to-particle partitioning
145 processes, amine-containing particles from both summer and winter were classified
146 into three types based on mass spectral patterns. The aminium sulfate and nitrate salt
147 formation processes and internal mixing state with ammonium were used to deduce
148 the relationship between amines and ammonium in the particle phase and the
149 influence of amines on particle acidity.

150 **2 Methods**

151 **2.1 Aerosol sampling**

152 Ambient single particles were collected and analyzed using a SPAMS at the
153 Guangdong Atmospheric Supersite (22.73° N, 112.93° E), a rural site in Heshan City
154 in the Pearl River Delta (PRD), China (Figure S1). The sampling site is surrounded by
155 villages and experiences little influence from local industrial emissions (Cheng et al.,
156 2017). The SPAMS was installed at the top of the main building, and aerosols were
157 introduced to the SPAMS through a 2.5 m copper tube. SPAMS sampling was
158 conducted continuously from 18 July to 1 August 2014 and from 27 January to 8
159 February 2015; several hours of data are missing due to technical maintenance.
160 During the sampling period, hourly O_3 concentrations were measured using an O_3

161 analyzer (model 49i, Thermo Scientific). Meteorological data, including temperature,
162 relative humidity, wind speed, and wind direction, were also measured during SPAMS
163 sampling.

164 **2.2 SPAMS**

165 SPAMS was designed by the Guangzhou Hexin Analytical Company based on
166 preexisting ATOFMS principles (Prather et al., 1994;Noble and Prather, 1996). The
167 setup and design of the SPAMS has been detailed previously (Li et al., 2011). Briefly,
168 single particles are sampled through an 80 μm critical orifice into the aerodynamic
169 lens at a flow rate of 75 ml min^{-1} . Then, the particles pass consecutively through two
170 laser beams (diode Nd:YAG, 532 nm) spaced 6 cm apart, and the aerodynamic
171 diameter of the single particle is calculated using the particle flight time and velocity
172 between the two laser beams. The single particle velocity is also used to calculate the
173 precise time at which to fire the desorption and ionization laser (Nd:YAG laser,
174 266nm), which is positioned 12 cm downstream from the second laser beam. After
175 ionization, the positive and negative ions are detected by a Z-shaped bipolar
176 time-of-flight mass spectrometer. In this work, the ionization laser pulse energy was
177 0.6 mJ and the power density was $1.06 \times 10^8 \text{ W cm}^{-2}$ throughout the campaign. The
178 size range of single particles detected by SPAMS ranged from 0.2 to 2 μm , calibrated
179 with standard polystyrene latex spheres (Nanosphere size standards, Duke Scientific
180 Corp., Palo Alto) of 0.22–2.0 μm diameter before and after the campaign (Cheng et al.,
181 2017).

182 **2.3 Data analysis**

183 Particle size and chemical composition were obtained via SPAMS mass spectral
184 analysis using the Computational Continuation Core (COCO; version 3.0) toolkit in
185 Matlab. According to the field studies of ATOFMS and SPAMS, it is difficult to
186 accurately determine the number concentration of ambient particles using SPAMS
187 alone due to the size-dependent transmission efficiencies of particles through
188 aerodynamic lens and composition dependent matrix effect (Gross et al., 2000;Pratt
189 and Prather, 2012). Thus, the particle counts and size distributions presented in this
190 work should be interpreted as semi-quantitative and serve as a basis of comparison

191 analysis (Healy et al., 2012). Based on previous studies using ATOFMS and SPAMS
192 instruments (Angelino et al., 2001;Huang et al., 2012;Zhang et al., 2012;Zauscher et
193 al., 2013;Healy et al., 2015), amine-containing particles were characterized by marker
194 ions, including m/z $^{59}(\text{CH}_3)_3\text{N}^+$, $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$, $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$, $^{101}(\text{C}_2\text{H}_5)_3\text{N}^+$,
195 $^{102}(\text{C}_3\text{H}_7)_2\text{NH}_2^+$, and $^{143}(\text{C}_3\text{H}_7)_3\text{N}^+$ (Table 1). In this work, a particle was identified as
196 amine-containing if it contained any of the marker ions listed above with a relative
197 peak area (defined as the percentage contribution of the target ion peak area to the
198 sum of all ion peak areas) greater than 1%. It should be noted that amine-containing
199 particles are operationally defined and not exclusive, which also contained various
200 chemical species in addition to amines. According to this criterion, 57452 and 68026
201 amine-containing particles were identified in summer and winter, respectively, which
202 accounted for 11.1 % and 9.4 % of the total detected particles. These number fractions
203 are consistent with previously reported observations in the PRD (Zhang et al., 2012).
204 However, due to the absence of fog events during the campaign, no dramatic increases
205 in amine-containing particles associated with high RH conditions ($\text{RH} > 90 \%$) were
206 observed. Marker ions of $^{59}(\text{CH}_3)_3\text{N}^+$, $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$, $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$ were detected
207 as the most abundant amines species during the sampling period, so particles
208 containing each marker ion were selected to investigate the possible sources and
209 characteristics of amine-containing particles. $^{30}\text{CH}_3\text{NH}^+$ is also an amine marker
210 which has been reported by other single particle studies (Phares et al.,
211 2003;Glagolenko and Phares, 2004). In this work the peak intensity of $^{30}\text{CH}_3\text{NH}^+$ was
212 much lower compared with other amine markers, and all the particles containing
213 $^{30}\text{CH}_3\text{NH}^+$ had strong signal of $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$, so the $^{30}\text{CH}_3\text{NH}^+$ -containing particles
214 were not specifically discussed. Ion of m/z +46 was detected in the ambient single
215 particles, which could be the amine marker of $^{46}(\text{CH}_3)_2\text{NH}_2^+$ and/or $^{46}\text{Na}_2^+$ according
216 to reported studies (Guazzotti et al., 2001;Gaston et al., 2011;Healy et al., 2015). In
217 this work the m/z +46-containing particles had no other amine markers as listed above,
218 besides, these particles were enriched with sodium salts like $^{62}\text{Na}_2\text{O}^+$, $^{81}\text{Na}_2\text{Cl}^+$ and
219 $^{147}\text{Na}(\text{NO}_3)_2^-$. Thus, m/z +46-containing particles were not classified as
220 amine-containing particles.

221 **3 Results and Discussion**

222 **3.1 Seasonal variation of amine-containing particles**

223 Spatial distributions of amine-containing particles associated with backward
224 trajectories (48 hour) of air masses at 500m levels above the ground during the
225 sampling period are shown in Figure 1. Cluster trajectories were calculated by
226 MeteoInfo (Wang, 2014), and the box plots were conducted by Igor Pro-based
227 program Histbox (Wu et al., 2018;Wu and Yu, 2018). In summer, high
228 amine-containing particle counts were associated with air masses of Cluster 3
229 (41.67%) and Cluster 4 (30.06%) (Figure 1a) from continent and South China Sea
230 separately, suggesting that the majority of amine-containing particles came from
231 anthropogenic and marine sources. However, in winter, large amounts of
232 amine-containing particles were associated with air masses of Cluster 4 (48.08%)
233 (Figure 1b), indicating that amine-containing particles were related primarily with
234 local emissions, such as animal husbandry, biomass burning, and vehicle exhaust.
235 Anthropogenic emissions from Foshan and Guangzhou may also have contributed, as
236 the sampling site is only 40 km and 56 km from these cities, respectively (Figure S1).

237 The amine-containing particle count observed in summer (57452) was lower than
238 it observed in winter (68026), but the abundance of amine-containing particles
239 relative to the total particle count was higher in summer (11.1%) than in winter (9.4%).
240 Temporal variations of total amine-containing particles and three amine marker ions
241 are shown in Figure 2. The increase of amine-containing particles was mostly
242 associated with high relative humidity (RH) at night in summer, while no direct
243 connection between particle counts and RH was found in winter (Figure S2 a and b).
244 High counts of amine-containing particles that extended in a few days were found
245 from 22 to 24 July (in summer) and from 5 to 8 February (in winter). Among the three
246 markers we considered, the most abundant amine marker was $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$, which
247 was detected in 90% and 86% of amine-containing particles in summer and winter
248 (Table 2), followed by $^{59}(\text{CH}_3)_3\text{N}^+$ and $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$ which were detected in less
249 than 10% of amine-containing particles during sampling period. The amine particles

250 containing $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$ and $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$ both exhibited similar variation
251 pattern with total amine-containing particles suggesting a similar emission source of
252 $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$ and $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$ (Figure 2). The temporal trend of
253 $^{59}(\text{CH}_3)_3\text{N}^+$ -containing particles were different from those of $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$ and
254 $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$; and the two sudden episodes of $^{59}(\text{CH}_3)_3\text{N}^+$ occurred from 27 to 29
255 July in summer were possibly due to the special emission sources of trimethylamine
256 (TMA).

257 The diurnal patterns of amine-containing particles are investigated in summer
258 and winter (Figure 3) and both showed higher count at night. The small increase from
259 6:00 to 9:00 LST throughout the campaign may have been due to local emissions
260 from vehicle exhaust (Cadle and Mulawa, 1980). Several field studies have revealed
261 the strong correlation between RH and particulate amines, suggesting that high RH in
262 fog events is favorable for the gas-to-particle partitioning of amines (Jeong et al.,
263 2011;Rehbein et al., 2011;Huang et al., 2012;Zhang et al., 2012). In this work,
264 although the increase of amine-containing particle count mostly occurred at night, no
265 obvious correlations between diurnal amine-containing particles and RH were found
266 in summer ($r^2=0.33$) and winter ($r^2=0.0003$) (Figure S2). The increase of
267 amine-containing particles at night may be influenced by particle acidity and emission
268 sources of amines (Murphy et al., 2007;Kurten et al., 2008;Silva et al., 2008).

269 **3.2 Characteristics of amine-containing particles**

270 The average mass spectra of amine-containing particles in summer and winter
271 are shown in Figure 4. The amine-containing particles were characterized by high
272 fractions of carbonaceous marker ions, including $^{27}\text{C}_2\text{H}_3^+$, $^{29}\text{C}_2\text{H}_5^+$, $^{36}\text{C}_3^+$, $^{37}\text{C}_3\text{H}^+$,
273 $^{43}\text{C}_2\text{H}_3\text{O}^+$, $^{48}\text{C}_4^+$, $^{51}\text{C}_4\text{H}_3^+$, $^{53}\text{C}_4\text{H}_5^+$, $^{60}\text{C}_5^+$, $^{63}\text{C}_5\text{H}_3^+$, $^{65}\text{C}_5\text{H}_5^+$, and $^{77}\text{C}_6\text{H}_5^+$; and amine
274 marker ions of $^{30}\text{CH}_3\text{NH}^+$, $^{59}(\text{CH}_3)_3\text{N}^+$, $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$ and $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$ in the
275 positive mass spectrum in both summer and winter. The negative mass spectrum was
276 characterized by strong carbon-nitrogen fragments like $^{26}\text{CN}^-$ and $^{42}\text{CNO}^-$, as well as
277 abundant secondary ions of $^{46}\text{NO}_2^-$, $^{62}\text{NO}_3^-$, $^{80}\text{SO}_3^-$, and $^{97}\text{HSO}_4^-$ in both summer and
278 winter. In many field studies, aged carbonaceous particles always contain abundant
279 secondary ions of sulfate, nitrate, and ammonium. Interestingly, in this work, the

280 signals of $^{18}\text{NH}_4^+$ were weak and only observed in less than 10% of amine-containing
281 particles in summer, but moderate signal of $^{18}\text{NH}_4^+$ was detected in half of
282 amine-containing particles in winter. The low $^{18}\text{NH}_4^+$ signal in amine-containing
283 particles may have been due to the emission sources of ammonia and particle acidity,
284 which will be discussed in Section 3.3.

285 The unscaled size-resolved number distributions of total amine-containing
286 particles and amine particles containing three marker ions of $^{59}(\text{CH}_3)_3\text{N}^+$,
287 $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$, and $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$ are shown in Figure 5. The amine-containing
288 particles exhibited unimodal distributions in the submicron mode from 0.4 to 1.5 μm
289 in both summer and winter, which may have resulted from gaseous amine
290 condensation on and/or reaction with fine mode particles from anthropogenic
291 emissions. Although amine-containing particles peaked at the size range of 0.5-0.7 μm
292 in both summer and winter, a broader size range of amine-containing particles was
293 observed in winter, which may be due to more complex anthropogenic emission
294 sources of primary particles in winter. The $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$ -containing particles showed
295 similar variation patterns as total amine-containing particles both in summer and
296 winter. However, $^{59}(\text{CH}_3)_3\text{N}^+$ - and $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$ -containing particles showed less
297 distinct peaks in winter.

298 **3.3 Mixing state and formation processes of amine-containing particles**

299 To investigate the aging state of amine-containing particles, the abundances of
300 sulfate-, nitrate-, and ammonium-containing amine particles are shown in Table 3.
301 More than 90% of amine-containing particles were found to be internally mixed with
302 sulfate throughout the sampling period. The abundance of nitrate in amine particles
303 increased from 43% in summer to 69% in winter. The high abundances of sulfate and
304 nitrate in amine-containing particles suggest the possible formation of aminium
305 sulfate and nitrate salts. Interestingly, only 8% of amine-containing particles mixed
306 with ammonium (NH_4^+) in summer, while the percentage increased dramatically to
307 54 % in winter, indicating a relatively NH_4^+ -poor state in summer and an NH_4^+ -rich
308 state in winter.

309 The seasonal differences of the mixing state of amines and NH_4^+ may be

310 influenced by the seasonal variation of source strength of NH_4^+ . To investigate the
311 temporal variation and abundance of NH_4^+ in total detected single particles, the total
312 NH_4^+ -containing particles were identified with relative area of $^{18}\text{NH}_4^+$ larger than 1%.
313 Using this criterion, 18336 and 235312 of NH_4^+ -containing particles were detected in
314 summer and winter separately, accounting for 3.6% and 32.6% of the total detected
315 particles. The averaged positive and negative ion mass spectra of NH_4^+ -containing
316 particles are exhibited in Figure 6. During entire sampling period the NH_4^+ -containing
317 particles were characterized by abundant hydrocarbon fragments and secondary
318 organic species like $^{43}\text{C}_2\text{H}_3\text{O}^+$ and $^{89}\text{HC}_2\text{O}_4^-$, as well as strong signals of $^{26}\text{CN}^-$,
319 $^{42}\text{CNO}^-$, $^{62}\text{NO}_3^-$ and $^{97}\text{HSO}_4^-$, indicating an aging state of NH_4^+ -containing particles.
320 Also, 20% of NH_4^+ -containing particles contained $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$, which indicates a
321 close connection between NH_3 and diethylamine (DEA), possibly due to the similar
322 emission sources.

323 Temporal variations of total amine-containing particles, total
324 ammonium-containing (NH_4^+ -containing) particles and particles containing both
325 ammonium and amine (NH_4^+ -amine) are shown in Figure 7. The total
326 NH_4^+ -containing particles and NH_4^+ -amine particles were both much lower in summer
327 than in winter. This seasonal difference may be due to the low emission sources of
328 ammonia and preferred partitioning in gas phase in summer. Backward trajectories
329 analysis (Figure 1) showed that in summer the air mass was mainly from south of the
330 sampling site and linked to the marine region with low emission of anthropogenic
331 pollutants. By contrast, in winter, the air mass was mainly from northwest of the
332 sampling site and associated with relatively polluted megacities like Guangzhou and
333 Foshan. RH does not seem to exert a major influence on particulate NH_4^+ (Huang et
334 al., 2012), because lower abundance of NH_4^+ was observed in summer ($\text{RH} = 72 \pm$
335 13%) than in winter ($\text{RH} = 63 \pm 11\%$).

336 The temporal variations of the peak areas of amines, ammonium, and the sum of
337 sulfate and nitrate in amine-containing particles are shown in Figure 8. The peak areas
338 of amines and the sum of nitrate and sulfate had similar variation patterns both in
339 summer and winter. The linear regression between them showed robust correlations

340 both in summer ($r^2=0.74$) and winter ($r^2=0.88$) (Figure S3), indicating the formation
341 of ammonium salts. Low peak area of ammonium was found in the amine-containing
342 particles in summer which was in accordance with the small amount of NH_4^+ -amine
343 particles. However, in winter, the peak area of ammonium was comparable with
344 amines and they both exhibited similar temporal trends. The sum of the sulfate and
345 nitrate peak areas had a higher increase rate than the amine peak area from 6 to 8
346 February, which may have been caused by an increase of ammonium during this
347 period. In this work the particle acidity of amine-containing particles is represented by
348 the relative acidity ratio (R_a), which is defined as the ratio of the sum of the sulfate
349 and nitrate peak areas divided by the ammonium peak area (Denkenberger et al.,
350 2007;Pratt et al., 2009;Cheng et al., 2017). Huang et al. (2013) obtained a robust
351 correlation ($r^2=0.82$) between the particle acidity calculated from inorganic ions
352 obtained from MARGA and relative acidity ratio obtained from single particle mass
353 spectrometer, allowing us to use R_a for comparison of particle acidity(Huang et al.,
354 2013). The R_a was 326 ± 326 in summer and 31 ± 13 in winter (Figure 8), indicating
355 that the amine-containing particles were more acidic in summer than in winter.

356 Although high acidity promotes gaseous ammonia partitioning, extremely low
357 ammonium peak areas were found for the amine-containing particles in summer
358 (Figure 8), which may be associated with ammonium–amine exchange reactions in
359 addition to the low emission source of ammonia. The exchange between amine gases
360 and particulate NH_3 and/or ammonium highly depends on the RH and particle acidity
361 (Chan and Chan, 2013;Chu and Chan, 2016). According to the study of Sauerwein
362 and Chan, the co-uptake of dimethylamine (DMA) and ammonia (NH_3) by sulfuric
363 acid particles at 50% RH led to particle-phase dimethylammonium (DMAH^+) to
364 ammonium (NH_4^+) molar ratio up to four times that of gas-phase DMA to ammonia
365 molar ratio (0.1 and 0.5), suggesting the displacement of NH_4^+ by DMA during the
366 uptake process (Sauerwein and Chan, 2017). In this work, the ambient RH and acidic
367 particles containing abundant sulfate and nitrate were similar to the experimental
368 conditions used in Sauerwein and Chan (2017). In summer 8% of amine-containing
369 particles contained NH_4^+ , while 25% of ammonium-containing particles contained

370 amines (Figure 7). Although the gas-phase concentrations of amines and NH_3 are not
371 quantified, higher abundance of amines in ammonium-containing particles than that of
372 ammonium in amine-containing particles suggests a possible ammonium–amine
373 exchange reactions in acidic particles in summer.

374 As strong bases, the presence of amines could have an impact on the particle
375 acidity. After including amines along with the ammonium in the relative acidity ratio
376 calculation, the new R_a' values (redefined as the ratio of the sum of the sulfate and
377 nitrate peak areas to the sum of the ammonium and amine peak areas) decrease to 11
378 ± 4 and 10 ± 2 in summer and winter, respectively, which are 30 and 3 times lower
379 than R_a values. Besides, R_a' showed no obvious seasonal change of particle acidity,
380 which suggests that amines could be a buffer for the particle acidity of
381 ammonium-poor particles, implying that it is reasonable to consider amines to
382 calculate particle acidity and actual pH. In addition, the presence of aminium salts
383 affects the water activities and osmotic coefficients of aqueous solutions, which may
384 influence the calculation of pH using aerosol thermodynamic models (Sauerwein et al.,
385 2015). Furthermore, it should be noted that the measured pH of bulk ambient aerosols
386 may not be representative of the actual single particle acidity. Hence, the mixing state
387 of aerosols should be considered in order to comprehensively estimate the aerosol pH
388 (Pratt et al., 2009). Several recent studies have reported a ‘missing’ source of sulfate
389 produced from the oxidation of SO_2 by NO_2 during haze episodes with high ambient
390 relative humidity in northern China, and the neutralization of particulate ammonium is
391 a key factor in this formation mechanism (Cheng et al., 2016; Wang et al., 2016). Our
392 study reveals that amines have a potential influence on particle acidity, which could
393 also impact this sulfate formation process during haze episodes. In order to discuss the
394 potential role of amines in this sulfate formation pathway, real-time concentrations of
395 amines, ammonium, sulfate, nitrate, and their precursors must be available. The
396 results of this study suggest that amine chemistry involving particle acidity and
397 mixing state with sulfate, nitrate and ammonium may have an important role in the
398 aging process of particles in regions with high concentration of amines.

399 **4 Summary and Conclusions**

400 Amine-containing particles were investigated using a single particle aerosol mass
401 spectrometer from 18 July to 1 August 2014, and from 27 January to 8 February 2015
402 in Heshan, China. Amine-containing particles accounted for 11.1 % and 9.4 % of the
403 total detected single particles in summer and winter, respectively; both seasons were
404 dominated by amine marker of $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$ in 90% and 86% of amine-containing
405 particles in summer and winter, respectively. Amine markers of $^{59}(\text{CH}_3)_3\text{N}^+$ and
406 $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$ were detected in 4.5% and 5.5% of amine-containing particles in
407 summer, while their percentages both increased two times in winter. The amine
408 particles contained $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$ and $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$ both exhibited similar
409 variation pattern with total amine-containing particles suggesting a similar emission
410 source of $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$ and $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$, while the $^{59}(\text{CH}_3)_3\text{N}^+$ -containing
411 particles showed different temporal trends, and two sudden increase episodes of
412 $^{59}(\text{CH}_3)_3\text{N}^+$ in summer was possibly due to the special sources of trimethylamine.
413 Although the increase of amine-containing particle count mostly occurred at night, no
414 obvious correlations between amine-containing particles and RH were found in
415 summer ($r^2=0.33$) and winter ($r^2=0.0003$). More than 90% of amine-containing
416 particles contained strong signals of sulfate throughout the sampling period, while 43%
417 and 69% of amine particles contained nitrate in summer in winter. Robust correlations
418 between the peak intensities of amines and the sum of nitrate and sulfate suggested
419 the possible formation of aminium sulfate and nitrate salts. Only 8% of amine
420 particles contained ammonium in summer, while the percentage increased
421 dramatically to 54% in winter. Due to the lower percentage of total
422 ammonium-containing particles in summer (3.6%) than it in winter (32.6%), the
423 relatively ammonium-poor state of amine-containing particles in summer may be due
424 to the lower abundance of ammonia/ammonium in gas and particle phase. Besides, 8%
425 of amine-containing particles contained ammonium while 25% of
426 ammonium-containing particles contained amines in summer, suggesting a possible
427 contribution of ammonium–amine exchange reactions to the low abundance of

428 ammonium in amine-containing particles at high ambient RH (72 ± 13 %) in summer.
429 In order to estimate the particle acidity, the relative acidity ratio (R_a), defined as the
430 ratio of the sum of the sulfate and nitrate peak areas divided by the ammonium peak
431 area, was calculated and showed higher values in summer (326 ± 326) than (31 ± 13)
432 in winter, suggesting the amine-containing particles were more acidic in summer than
433 in winter. However, after including amines along with the ammonium in the acidity
434 calculation, the new R_a' values showed no distinct seasonal change (summer: 11 ± 4 ;
435 winter: 10 ± 2), suggesting that it is reasonable to consider amines when estimating
436 particle acidity.

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438

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441 manuscript with contributions from all co-authors.

442

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445

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461 **References**

- 462 Angelino, S., Suess, D. T., and Prather, K. A.: Formation of aerosol particles from
463 reactions of secondary and tertiary alkylamines: Characterization by aerosol
464 time-of-flight mass spectrometry, *Environmental Science & Technology*, 35,
465 3130-3138, Doi 10.1021/Es0015444, 2001.
- 466 Berndt, T., Stratmann, F., Sipila, M., Vanhanen, J., Petaja, T., Mikkila, J., Gruner, A.,
467 Spindler, G., Mauldin, R. L., Curtius, J., Kulmala, M., and Heintzenberg, J.:
468 Laboratory study on new particle formation from the reaction OH + SO₂:
469 influence of experimental conditions, H₂O vapour, NH₃ and the amine
470 tert-butylamine on the overall process, *Atmospheric Chemistry and Physics*, 10,
471 7101-7116, 10.5194/acp-10-7101-2010, 2010.
- 472 Bzdek, B., Ridge, D., and Johnston, M.: Amine exchange into ammonium bisulfate
473 and ammonium nitrate nuclei, *Atmospheric Chemistry and Physics*, 10,
474 3495-3503, 2010.
- 475 Cadle, S. H., and Mulawa, P. A.: Low-molecular-weight aliphatic amines in exhaust
476 from catalyst-equipped cars, *Environmental science & technology*, 14, 718-723,
477 1980.
- 478 Calderon, S. M., Poor, N. D., and Campbell, S. W.: Estimation of the particle and gas
479 scavenging contributions to wet deposition of organic nitrogen, *Atmospheric*
480 *Environment*, 41, 4281-4290, 10.1016/j.atmosenv.2006.06.067, 2007.
- 481 Chan, L. P., and Chan, C. K.: Role of the Aerosol Phase State in Ammonia/Amines
482 Exchange Reactions, *Environmental Science & Technology*, 47, 5755-5762,
483 10.1021/es4004685, 2013.
- 484 Cheng, C., Li, M., Chan, C. K., Tong, H., Chen, C., Chen, D., Wu, D., Li, L., Wu, C.,
485 Cheng, P., Gao, W., Huang, Z., Li, X., Zhang, Z., Fu, Z., Bi, Y., and Zhou, Z.:
486 Mixing state of oxalic acid containing particles in the rural area of Pearl River
487 Delta, China: implications for the formation mechanism of oxalic acid,
488 *Atmospheric Chemistry and Physics*, 17, 9519-9533, 10.5194/acp-17-9519-2017,
489 2017.
- 490 Cheng, Y., Zheng, G., Wei, C., Mu, Q., Zheng, B., Wang, Z., Gao, M., Zhang, Q., He,
491 K., Carmichael, G., Pöschl, U., and Su, H.: Reactive nitrogen chemistry in
492 aerosol water as a source of sulfate during haze events in China, *Science*
493 *Advances*, 2, 10.1126/sciadv.1601530, 2016.
- 494 Chu, Y., and Chan, C. K.: Reactive Uptake of Dimethylamine by Ammonium Sulfate
495 and Ammonium Sulfate–Sucrose Mixed Particles, *The Journal of Physical*
496 *Chemistry A*, 121, 206-215, 10.1021/acs.jpca.6b10692, 2016.

497 Chu, Y., and Chan, C. K.: Role of oleic acid coating in the heterogeneous uptake of
498 dimethylamine by ammonium sulfate particles, *Aerosol Science and Technology*,
499 51, 988-997, 10.1080/02786826.2017.1323072, 2017.

500 Chu, Y. X., Sauerwein, M., and Chan, C. K.: Hygroscopic and phase transition
501 properties of alkyl aminium sulfates at low relative humidities, *Phys Chem Chem*
502 *Phys*, 17, 19789-19796, 10.1039/c5cp02404h, 2015.

503 Day, D. A., Takahama, S., Gilardoni, S., and Russell, L. M.: Organic composition of
504 single and submicron particles in different regions of western North America and
505 the eastern Pacific during INTEX-B 2006, *Atmospheric Chemistry and Physics*,
506 9, 5433-5446, 2009.

507 Denkenberger, K. A., Moffet, R. C., Holecek, J. C., Rebotier, T. P., and Prather, K. A.:
508 Real-time, single-particle measurements of oligomers in aged ambient aerosol
509 particles, *Environmental Science & Technology*, 41, 5439-5446,
510 10.1021/es070329l, 2007.

511 Gaston, C. J., Furutani, H., Guazzotti, S. A., Coffee, K. R., Bates, T. S., Quinn, P. K.,
512 Aluwihare, L. I., Mitchell, B. G., and Prather, K. A.: Unique ocean - derived
513 particles serve as a proxy for changes in ocean chemistry, *Journal of Geophysical*
514 *Research: Atmospheres* (1984 - 2012), 116, 10.1029/2010JD015289, 2011.

515 Ge, X. L., Wexler, A. S., and Clegg, S. L.: Atmospheric amines - Part II.
516 Thermodynamic properties and gas/particle partitioning, *Atmospheric*
517 *Environment*, 45, 561-577, 10.1016/j.atmosenv.2010.10.013, 2011a.

518 Ge, X. L., Wexler, A. S., and Clegg, S. L.: Atmospheric amines - Part I. A review,
519 *Atmospheric Environment*, 45, 524-546, DOI 10.1016/j.atmosenv.2010.10.012,
520 2011b.

521 Gilardoni, S., Liu, S., Takahama, S., Russell, L. M., Allan, J. D., Steinbrecher, R.,
522 Jimenez, J. L., De Carlo, P. F., Dunlea, E. J., and Baumgardner, D.:
523 Characterization of organic ambient aerosol during MIRAGE 2006 on three
524 platforms, *Atmospheric Chemistry and Physics*, 9, 5417-5432, 2009.

525 Glagolenko, S., and Phares, D. J.: Single-particle analysis of ultrafine aerosol in
526 College Station, Texas, *Journal of Geophysical Research-Atmospheres*, 109,
527 10.1029/2004jd004621, 2004.

528 Gross, D. S., Galli, M. E., Silva, P. J., and Prather, K. A.: Relative sensitivity factors
529 for alkali metal and ammonium cations in single particle aerosol time-of-flight
530 mass spectra, *Anal. Chem.*, 72, 416-422, Doi 10.1021/Ac990434g, 2000.

531 Guazzotti, S. A., Whiteaker, J. R., Suess, D., Coffee, K. R., and Prather, K. A.:
532 Real-time measurements of the chemical composition of size-resolved particles
533 during a Santa Ana wind episode, California USA, *Atmospheric Environment*, 35,
534 3229-3240, 2001.

535 Healy, R., Sciare, J., Poulain, L., Kamili, K., Merkel, M., Müller, T., Wiedensohler, A.,
536 Eckhardt, S., Stohl, A., and Sarda-Estève, R.: Sources and mixing state of
537 size-resolved elemental carbon particles in a European megacity: Paris,
538 *Atmospheric Chemistry and Physics*, 12, 1681-1700, 2012.

539 Healy, R. M., Evans, G. J., Murphy, M., Sierau, B., Arndt, J., McGillicuddy, E.,
540 O'Connor, I. P., Sodeau, J. R., and Wenger, J. C.: Single-particle speciation of

541 alkylamines in ambient aerosol at five European sites, *Analytical and*
542 *Bioanalytical Chemistry*, 407, 5899-5909, 10.1007/s00216-014-8092-1, 2015.

543 Huang, Y., Chen, H., Wang, L., Yang, X., and Chen, J.: Single particle analysis of
544 amines in ambient aerosol in Shanghai, *Environmental Chemistry*, 9, 202-210,
545 10.1071/EN11145, 2012.

546 Huang, Y., Li, L., Li, J., Wang, X., Chen, H., Chen, J., Yang, X., Gross, D. S., Wang,
547 H., Qiao, L., and Chen, C.: A case study of the highly time-resolved evolution of
548 aerosol chemical and optical properties in urban Shanghai, China, *Atmospheric*
549 *Chemistry and Physics*, 13, 3931-3944, 10.5194/acp-13-3931-2013, 2013.

550 Jeong, C. H., McGuire, M. L., Godri, K. J., Slowik, J. G., Rehbein, P. J. G., and Evans,
551 G. J.: Quantification of aerosol chemical composition using continuous single
552 particle measurements, *Atmospheric Chemistry and Physics*, 11, 7027-7044,
553 10.5194/acp-11-7027-2011, 2011.

554 Kurten, T., Loukonen, V., Vehkamäki, H., and Kulmala, M.: Amines are likely to
555 enhance neutral and ion-induced sulfuric acid-water nucleation in the atmosphere
556 more effectively than ammonia, *Atmospheric Chemistry and Physics*, 8,
557 4095-4103, 2008.

558 Lee, D., and Wexler, A. S.: Atmospheric amines - Part III: Photochemistry and toxicity,
559 *Atmospheric Environment*, 71, 95-103, DOI 10.1016/j.atmosenv.2013.01.058,
560 2013.

561 Li, L., Huang, Z. X., Dong, J. G., Li, M., Gao, W., Nian, H. Q., Fu, Z., Zhang, G. H.,
562 Bi, X. H., Cheng, P., and Zhou, Z.: Real time bipolar time-of-flight mass
563 spectrometer for analyzing single aerosol particles, *Int J Mass Spectrom*, 303,
564 118-124, 10.1016/j.ijms.2011.01.017, 2011.

565 Liu, S., Takahama, S., Russell, L. M., Gilardoni, S., and Baumgardner, D.:
566 Oxygenated organic functional groups and their sources in single and submicron
567 organic particles in MILAGRO 2006 campaign, *Atmospheric Chemistry and*
568 *Physics*, 9, 6849-6863, 2009.

569 Liu, Y., Han, C., Liu, C., Ma, J., Ma, Q., and He, H.: Differences in the reactivity of
570 ammonium salts with methylamine, *Atmospheric Chemistry and Physics*, 12,
571 4855-4865, 2012.

572 Lloyd, J. A., Heaton, K. J., and Johnston, M. V.: Reactive uptake of trimethylamine
573 into ammonium nitrate particles, *The Journal of Physical Chemistry A*, 113,
574 4840-4843, 10.1021/jp900634d, 2009.

575 Malloy, Q. G. J., Li, Q., Warren, B., Cocker III, D. R., Erupe, M. E., and Silva, P. J.:
576 Secondary organic aerosol formation from primary aliphatic amines with
577 NO₃ radical, *Atmos. Chem. Phys.*, 9, 2051-2060,
578 10.5194/acp-9-2051-2009, 2009.

579 Moffet, R. C., de Foy, B., Molina, L. T., Molina, M. J., and Prather, K. A.:
580 Measurement of ambient aerosols in northern Mexico City by single particle
581 mass spectrometry, *Atmospheric Chemistry and Physics*, 8, 4499-4516,
582 10.5194/acp-8-4499-2008, 2008.

583 Murphy, S. M., Sorooshian, A., Kroll, J. H., Ng, N. L., Chhabra, P., Tong, C., Surratt,
584 J. D., Knipping, E., Flagan, R. C., and Seinfeld, J. H.: Secondary aerosol

585 formation from atmospheric reactions of aliphatic amines, *Atmospheric*
586 *Chemistry and Physics*, 7, 2313-2337, 2007.

587 Ngwabie, N. M., Schade, G. W., Custer, T. G., Linke, S., and Hinz, T.: Volatile organic
588 compound emission and other trace gases from selected animal buildings,
589 *Landbauforsch Volk*, 57, 273-284, 2007.

590 Noble, C. A., and Prather, K. A.: Real-time measurement of correlated size and
591 composition profiles of individual atmospheric aerosol particles, *Environmental*
592 *science & technology*, 30, 2667-2680, 1996.

593 Phares, D. J., Rhoads, K. P., Johnston, M. V., and Wexler, A. S.: Size-resolved
594 ultrafine particle composition analysis - 2. Houston, *Journal of Geophysical*
595 *Research-Atmospheres*, 108, 10.1029/2001jd001212, 2003.

596 Place, P. F., Ziemba, L. D., and Griffin, R. J.: Observations of nucleation-mode
597 particle events and size distributions at a rural New England site, *Atmospheric*
598 *Environment*, 44, 88-94, 10.1016/j.atmosenv.2009.09.030, 2010.

599 Prather, K. A., Nordmeyer, T., and Salt, K.: Real-time characterization of individual
600 aerosol particles using time-of-flight mass spectrometry, *Anal. Chem.*, 66,
601 1403-1407, 1994.

602 Pratt, K. A., Hatch, L. E., and Prather, K. A.: Seasonal Volatility Dependence of
603 Ambient Particle Phase Amines, *Environmental Science & Technology*, 43,
604 5276-5281, 10.1021/es803189n, 2009.

605 Pratt, K. A., and Prather, K. A.: Mass spectrometry of atmospheric aerosols—Recent
606 developments and applications. Part II: On - line mass spectrometry techniques,
607 *Mass Spectrom Rev*, 31, 17-48, 2012.

608 Qiu, C., Wang, L., Lal, V., Khalizov, A. F., and Zhang, R.: Heterogeneous reactions of
609 alkylamines with ammonium sulfate and ammonium bisulfate, *Environmental*
610 *science & technology*, 45, 4748-4755, 2011.

611 Rappert, S., and Muller, R.: Odor compounds in waste gas emissions from agricultural
612 operations and food industries, *Waste Manage*, 25, 887-907,
613 10.1016/j.wasman.2005.07.008, 2005.

614 Rehbein, P. J. G., Jeong, C. H., McGuire, M. L., Yao, X. H., Corbin, J. C., and Evans,
615 G. J.: Cloud and Fog Processing Enhanced Gas-to-Particle Partitioning of
616 Trimethylamine, *Environmental Science & Technology*, 45, 4346-4352,
617 10.1021/es1042113, 2011.

618 Rogge, W. F., Hildemann, L. M., Mazurek, M. A., and Cass, G. R.: Sources of Fine
619 Organic Aerosol .6. Cigarette-Smoke in the Urban Atmosphere, *Environmental*
620 *Science & Technology*, 28, 1375-1388, 10.1021/Es00056a030, 1994.

621 Russell, L. M., Takahama, S., Liu, S., Hawkins, L. N., Covert, D. S., Quinn, P. K., and
622 Bates, T. S.: Oxygenated fraction and mass of organic aerosol from direct
623 emission and atmospheric processing measured on the R/V Ronald Brown during
624 TEXAQS/GoMACCS 2006, *Journal of Geophysical Research-Atmospheres*, 114,
625 10.1029/2008jd011275, 2009.

626 Sauerwein, M., Clegg, S. L., and Chan, C. K.: Water Activities and Osmotic
627 Coefficients of Aqueous Solutions of Five Alkylammonium Sulfates and Their
628 Mixtures with H₂SO₄ at 25(o)C, *Aerosol Science and Technology*, 49, 566-579,

629 10.1080/02786826.2015.1043045, 2015.

630 Sauerwein, M., and Chan, C. K.: Heterogeneous uptake of ammonia and
631 dimethylamine into sulfuric and oxalic acid particles, *Atmos. Chem. Phys.*, 17,
632 6323-6339, 10.5194/acp-17-6323-2017, 2017.

633 Silva, P. J., Erupe, M. E., Price, D., Elias, J., Malloy, Q. G. J., Li, Q., Warren, B., and
634 Cocker, D. R.: Trimethylamine as precursor to secondary organic aerosol
635 formation via nitrate radical reaction in the atmosphere, *Environmental Science
636 & Technology*, 42, 4689-4696, Doi 10.1021/Es703016v, 2008.

637 Sorooshian, A., Ng, N. L., Chan, A. W. H., Feingold, G., Flagan, R. C., and Seinfeld, J.
638 H.: Particulate organic acids and overall water-soluble aerosol composition
639 measurements from the 2006 Gulf of Mexico Atmospheric Composition and
640 Climate Study (GoMACCS), *Journal of Geophysical Research-Atmospheres*,
641 112, 10.1029/2007jd008537, 2007.

642 Sorooshian, A., Murphy, S. N., Hersey, S., Gates, H., Padro, L. T., Nenes, A., Brechtel,
643 F. J., Jonsson, H., Flagan, R. C., and Seinfeld, J. H.: Comprehensive airborne
644 characterization of aerosol from a major bovine source, *Atmospheric Chemistry
645 and Physics*, 8, 5489-5520, 2008.

646 Wang, G., Zhang, R., Gomez, M. E., Yang, L., Levy Zamora, M., Hu, M., Lin, Y.,
647 Peng, J., Guo, S., Meng, J., Li, J., Cheng, C., Hu, T., Ren, Y., Wang, Y., Gao, J.,
648 Cao, J., An, Z., Zhou, W., Li, G., Wang, J., Tian, P., Marrero-Ortiz, W., Secret, J.,
649 Du, Z., Zheng, J., Shang, D., Zeng, L., Shao, M., Wang, W., Huang, Y., Wang, Y.,
650 Zhu, Y., Li, Y., Hu, J., Pan, B., Cai, L., Cheng, Y., Ji, Y., Zhang, F., Rosenfeld, D.,
651 Liss, P. S., Duce, R. A., Kolb, C. E., and Molina, M. J.: Persistent sulfate
652 formation from London Fog to Chinese haze, *Proceedings of the National
653 Academy of Sciences*, 113, 13630-13635, 10.1073/pnas.1616540113, 2016.

654 Wang, L., Khalizov, A. F., Zheng, J., Xu, W., Ma, Y., Lal, V., and Zhang, R. Y.:
655 Atmospheric nanoparticles formed from heterogeneous reactions of organics, *Nat
656 Geosci*, 3, 238-242, 10.1038/NNGEO778, 2010.

657 Wang, Y. Q.: MeteoInfo: GIS software for meteorological data visualization and
658 analysis, *Meteorol Appl*, 21, 360-368, 10.1002/met.1345, 2014.

659 Williams, B. J., Goldstein, A. H., Kreisberg, N. M., Hering, S. V., Worsnop, D. R.,
660 Ulbrich, I. M., Docherty, K. S., and Jimenez, J. L.: Major components of
661 atmospheric organic aerosol in southern California as determined by hourly
662 measurements of source marker compounds, *Atmospheric Chemistry and Physics*,
663 10, 11577-11603, 10.5194/acp-10-11577-2010, 2010.

664 Wu, C., Wu, D., and Yu, J. Z.: Quantifying black carbon light absorption enhancement
665 with a novel statistical approach, *Atmospheric Chemistry and Physics*, 18,
666 289-309, 10.5194/acp-18-289-2018, 2018.

667 Wu, C., and Yu, J. Z.: Evaluation of linear regression techniques for atmospheric
668 applications: the importance of appropriate weighting, *Atmos Meas Tech*, 11,
669 1233-1250, 10.5194/amt-11-1233-2018, 2018.

670 Zauscher, M. D., Wang, Y., Moore, M. J. K., Gaston, C. J., and Prather, K. A.: Air
671 Quality Impact and Physicochemical Aging of Biomass Burning Aerosols during
672 the 2007 San Diego Wildfires, *Environmental Science & Technology*, 47,

673 7633-7643, 10.1021/es4004137, 2013.
674 Zhang, G., Bi, X., Chan, L. Y., Li, L., Wang, X., Feng, J., Sheng, G., Fu, J., Li, M.,
675 and Zhou, Z.: Enhanced trimethylamine-containing particles during fog events
676 detected by single particle aerosol mass spectrometry in urban Guangzhou, China,
677 Atmospheric Environment, 55, 121-126, 2012.

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717 **Tables and Figures**

718 **Table list:**

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720

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723

724 Table 3. The abundances of ammonium-, nitrate- and sulfate-containing amine
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726 **Figure captions:**

727 Figure 1. Spatial distributions of amine-containing particle counts associated with
728 backward trajectories (48 hour) of air masses at 500m levels above the ground during
729 the sampling period: (a) summer (from July 18 to August 1, 2014), (b) winter (from
730 January 27 to February 8, 2015).

731

732 Figure 2. Temporal variations of relative humidity (RH), temperature (T), total
733 amine-containing particles, and three major marker ions-containing amine particles
734 ($^{59}(\text{CH}_3)_3\text{N}^+$, $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$, $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$) in Heshan, China during sampling
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738 in Heshan.

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741 winter. The color bars represent each peak area corresponding to a specific ion in
742 individual particles.

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746 $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$, and $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$ in summer and winter in Heshan.

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749 summer and winter. The color bars represent each peak area corresponding to a
750 specific fraction in individual particles.

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753 ammonium-containing particles, and particles containing both ammonium and amine
754 (NH_4^+ -amine) during sampling period in Heshan.

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756 Figure 8. Temporal variations in the peak areas of amines, ammonium, and the sum of
757 sulfate and nitrate in amine-containing particles during summer and winter. The

758 relative acidity ratio (R_a), which was calculated as the ratio of the total sulfate and
759 nitrate peak areas to the ammonium peak area, is plotted as $\log(R_a)$.

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802 **Tables:**

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Table 1. Marker ions chosen for the amine-containing particles

Marker ion	Alkylamine assignment
$^{59}(\text{CH}_3)_3\text{N}^+$	Trimethylamine (TMA)
$^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$	Diethylamine (DEA)
$^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$	DEA, TEA, DPA
$^{101}(\text{C}_2\text{H}_5)_3\text{N}^+$	Triethylamine (TEA)
$^{102}(\text{C}_3\text{H}_7)_2\text{NH}_2^+$	Dipropylamine (DPA)
$^{143}(\text{C}_3\text{H}_7)_3\text{N}^+$	Tripropylamine (TPA)

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Table 2. Seasonal distributions of amine-containing particles and three major amine markers in summer and winter in the PRD, China.

Particle type	Summer (18/7-1/8, 2014)		Winter (27/1-8/2, 2015)	
	Count	Percentage (%) ^a	Count	Percentage (%) ^a
Total Amines	57452		68026	
$^{59}(\text{CH}_3)_3\text{N}^+$	2581	4.5	6894	10
$^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$	51442	90	58272	86
$^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$	3185	5.5	6119	9

^aThe percentage of each amine marker ion in total detected amine-containing particles.

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Table 3. The abundances of ammonium-, nitrate- and sulfate-containing amine particles in total amine-containing particles.

Marker ions	Summer	Winter
$^{18}\text{NH}_4^+$	8%	54%
$^{62}\text{NO}_3^-$	43%	69%
$^{97}\text{HSO}_4^-$	91%	94%

The marker ions of $^{18}\text{NH}_4^+$, $^{62}\text{NO}_3^-$ and $^{97}\text{HSO}_4^-$ were chosen to represent ammonium, nitrate and sulfate.

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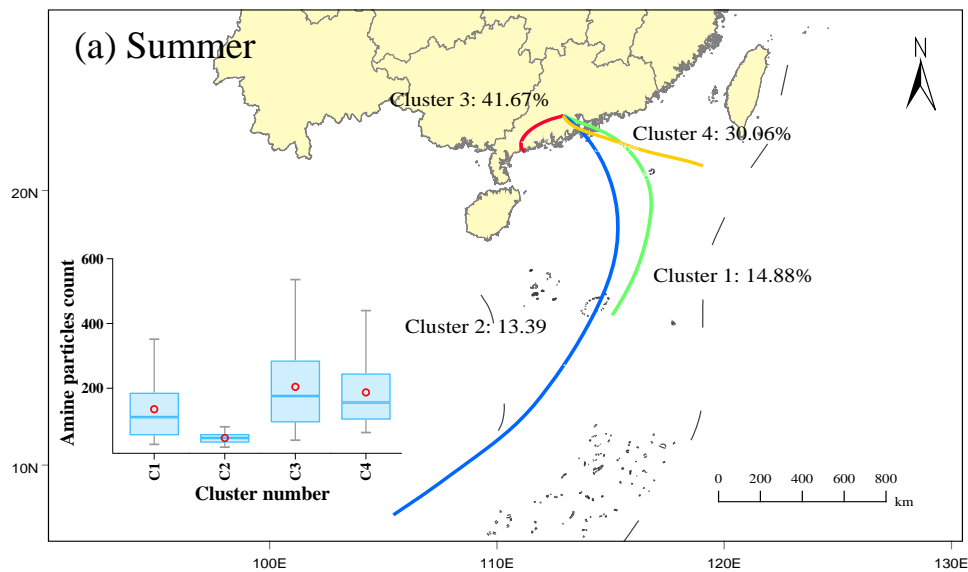
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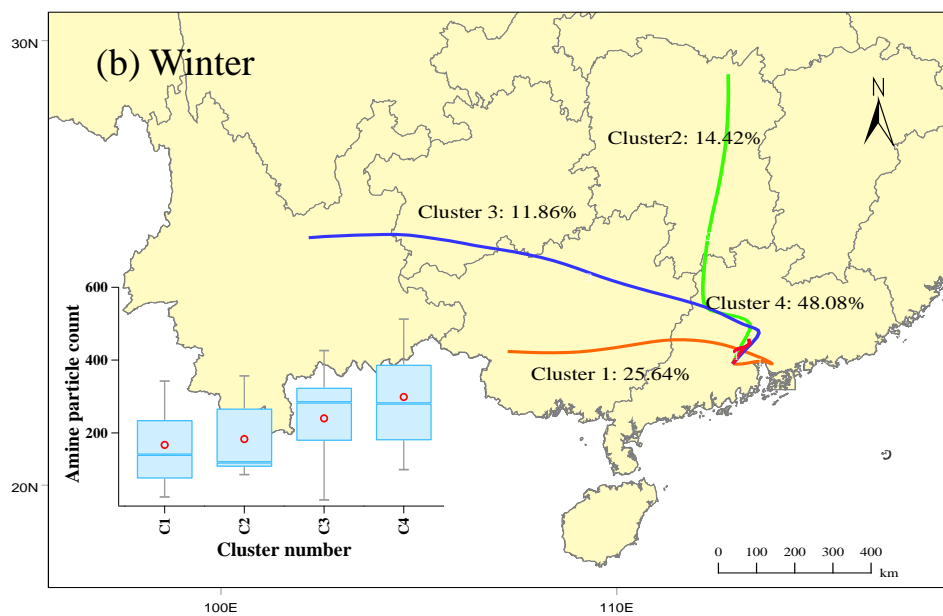
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820 **Figures:**

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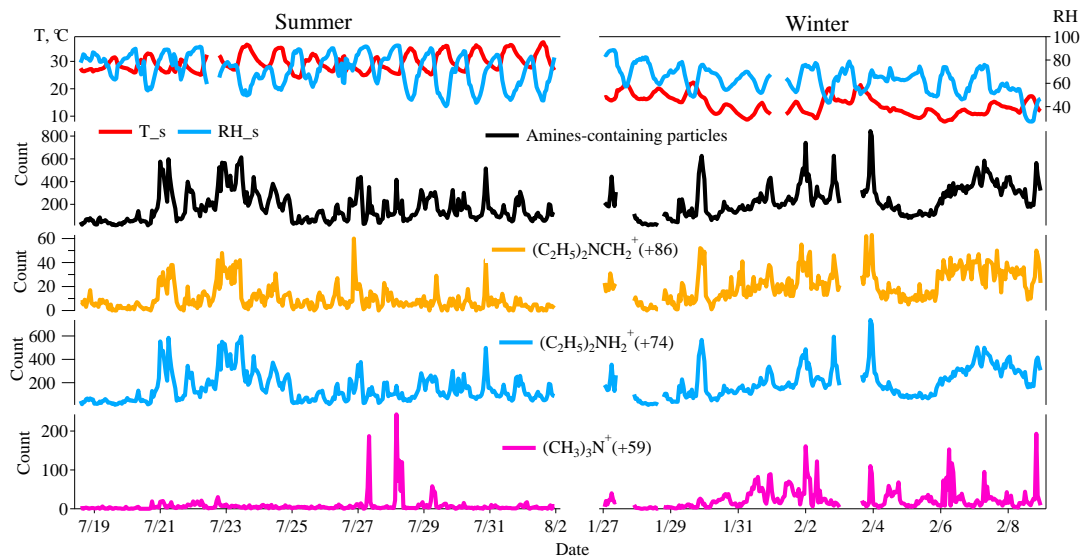
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824 Figure 1. Spatial distributions of amine-containing particle counts associated with
825 backward trajectories (48 hour) of air masses at 500m levels above the ground during
826 the sampling period: (a) summer (from July 18 to August 1, 2014), (b) winter (from
827 January 27 to February 8, 2015).

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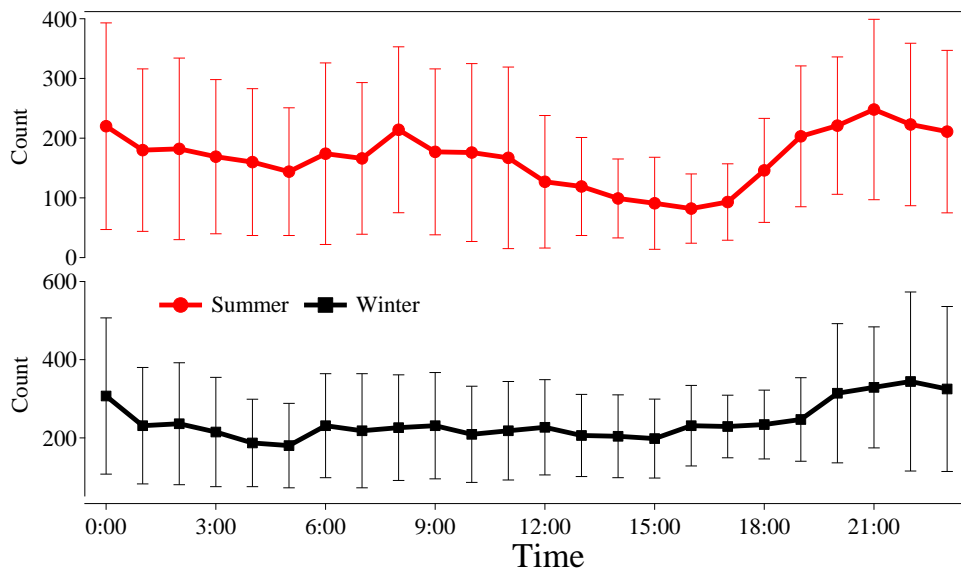
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831 Figure 2. Temporal variations of relative humidity (RH), temperature (T), total
 832 amine-containing particles, and three major marker ions-containing amine particles
 833 ($^{59}\text{(CH}_3)_3\text{N}^+$, $^{74}\text{(C}_2\text{H}_5)_2\text{NH}_2^+$, $^{86}\text{(C}_2\text{H}_5)_2\text{NCH}_2^+$) in Heshan, China during sampling periods.

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838 Figure 3. Diurnal variations of amine-containing particle counts in summer and winter
 839 in Heshan.

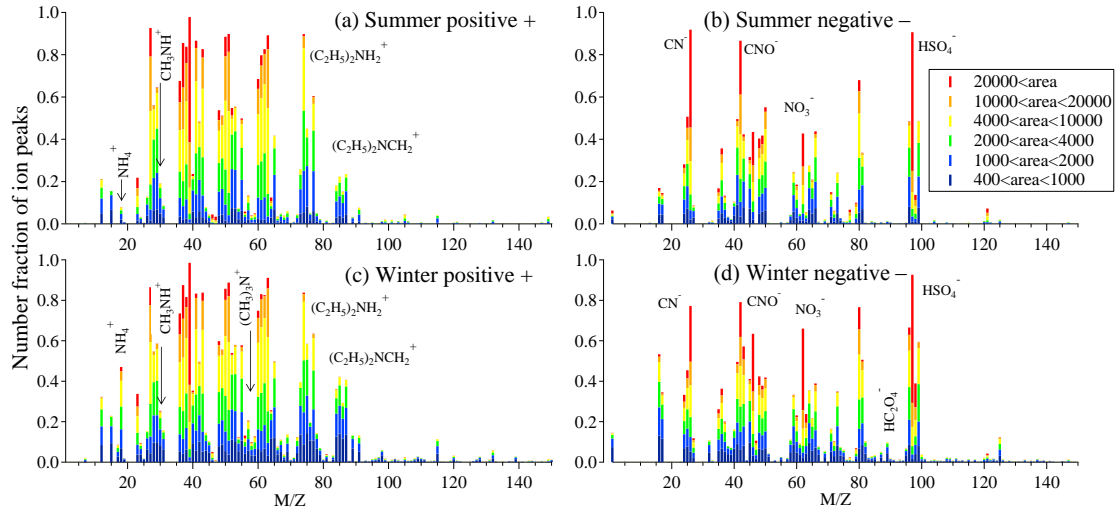
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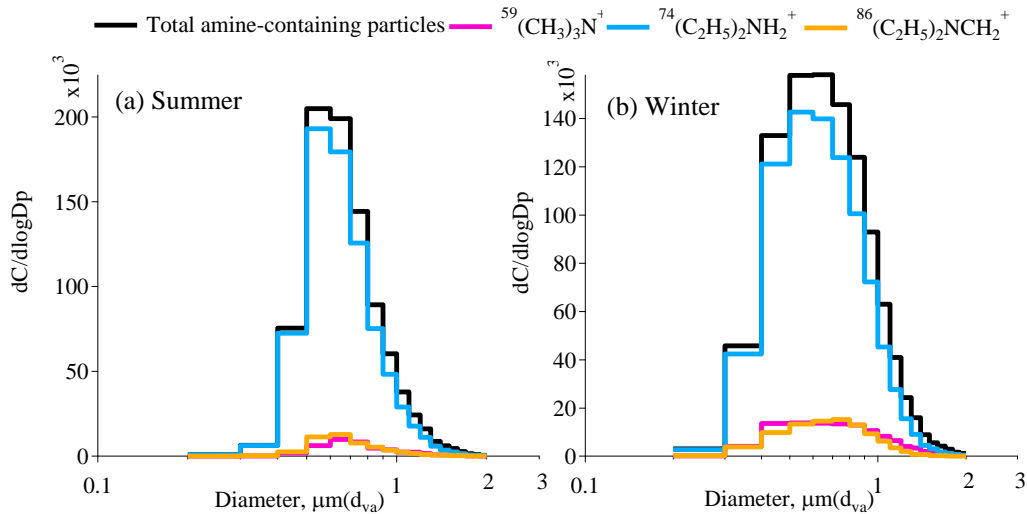
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Figure 4. Average ion mass spectra of amine-containing particles in summer and winter. The color bars represent each peak area corresponding to a specific ion in individual particles.

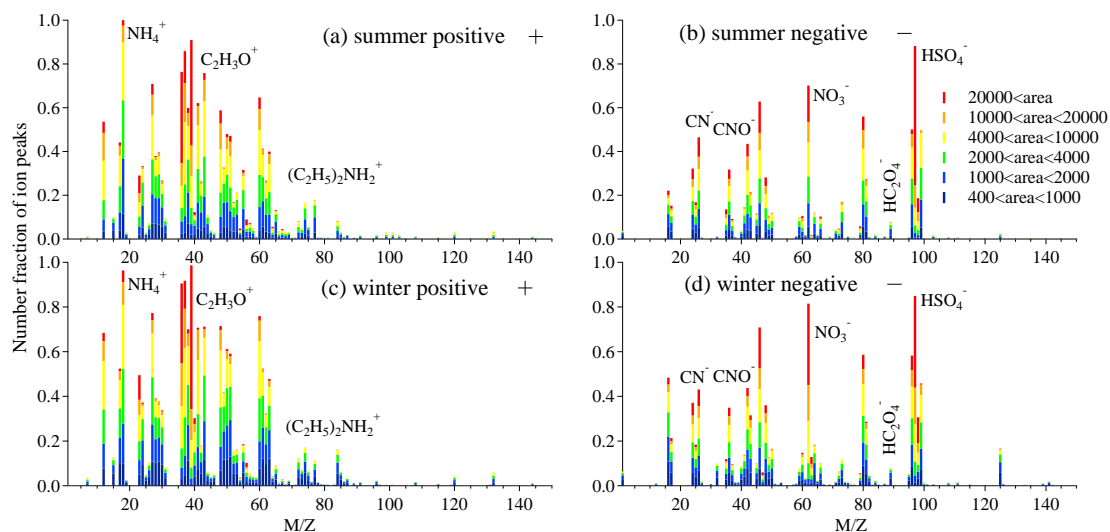


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Figure 5. Unscaled size-resolved number distributions of total amine-containing particles and amine particles containing three marker ions of $^{59}(\text{CH}_3)_3\text{N}^+$, $^{74}(\text{C}_2\text{H}_5)_2\text{NH}_2^+$, and $^{86}(\text{C}_2\text{H}_5)_2\text{NCH}_2^+$ in summer and winter in Heshan.

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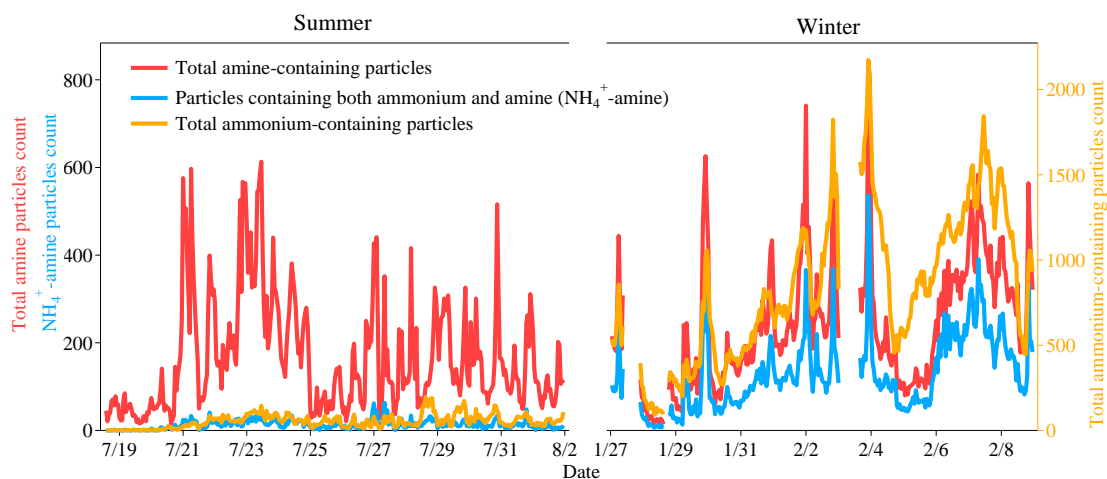
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Figure 6. Mass spectra of total ammonium-containing (NH_4^+ -containing) particles in summer and winter. The color bars represent each peak area corresponding to a specific fraction in individual particles.



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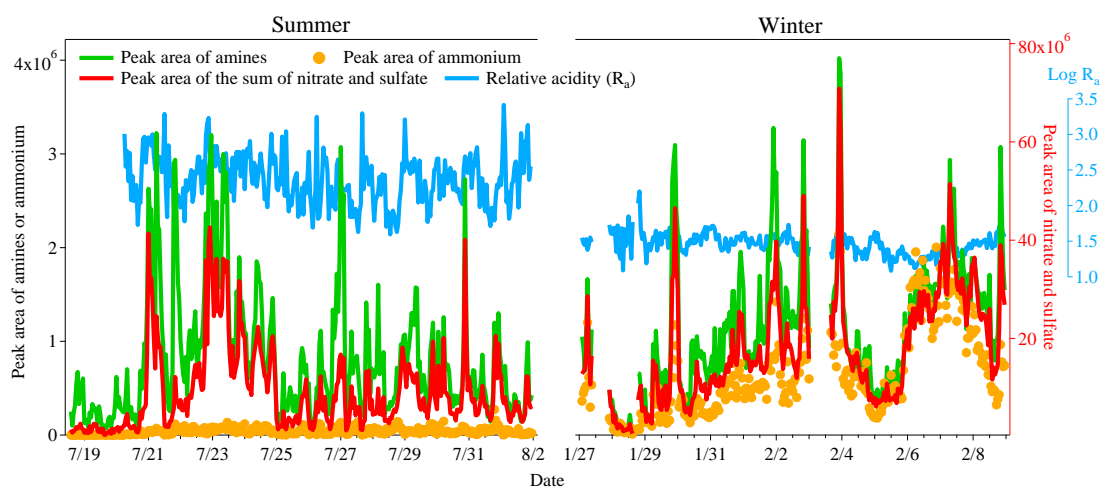
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Figure 7. Temporal variations of total amine-containing particles, total ammonium-containing particles and particles containing both ammonium and amine (NH_4^+ -amine) during sampling period in Heshan.

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887 Figure 8. Temporal variations in the peak areas of amines, ammonium, and the sum of
888 sulfate and nitrate in amine-containing particles during summer and winter. The
889 relative acidity ratio (R_a), which was calculated as the ratio of the total sulfate and
890 nitrate peak areas to the ammonium peak area, is plotted as $\log(R_a)$.

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