



Measurements of NO and NO₂ exchange between the atmosphere and *Quercus agrifolia*

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Abstract. NO₂ foliar deposition through the stomata of leaves has been identified as a significant sink of NO_x within a forest canopy. In this study, we investigated NO₂ and NO exchange between the atmosphere and the leaves of the native California oak tree *Quercus agrifolia* using a branch enclosure system. NO₂ detection was performed with laser-induced fluorescence (LIF), which excludes biases from other reactive nitrogen compounds and has a low detection limit of 5–50 ppt. We performed both light and dark experiments with concentrations between 0.5–10 ppb NO₂ and NO under constant ambient conditions. Deposition velocities for NO₂ during light and dark experiments were 0.123 ± 0.007 cm s⁻¹ and 0.015 ± 0.001 cm s⁻¹, respectively. Much slower deposition was seen for NO, with deposition velocities of 0.012 ± 0.002 cm s⁻¹ and 0.005 ± 0.002 cm s⁻¹ measured during light and dark experiments, respectively. This corresponded to a summed resistance of the stomata and mesophyll of 6.9 ± 0.9 s cm⁻¹ for NO₂ and 140 ± 40 s cm⁻¹ for NO. No significant compensation point was detected for NO₂ uptake, but compensation points ranging from 0.74–3.8 ppb were observed for NO. NO₂ and NO deposition velocities reported here are comparable both with previous leaf-level chamber studies and inferences from canopy-level field measurements. In parallel with these laboratory experiments, we have constructed a detailed 1-D atmospheric model to assess the contribution of leaf-level NO_x deposition to the total NO_x loss and NO_x canopy fluxes. Using the leaf uptake rates measured in the laboratory, these modeling studies suggest loss of NO_x to deposition in a California oak woodland competes with the pathways of HNO₃ and RONO₂ formation, with deposition making up 3–22% of the total NO_x loss. Additionally, foliar uptake of NO_x at these rates could account for ~15–30% canopy reduction of soil NO_x emissions.

25 1 Introduction

Nitrogen oxides (NO_x ≡ NO + NO₂) are a group of highly reactive trace gases that control the oxidative capacity of the atmosphere, regulating the amounts of ozone, hydroxyl radicals, volatile organic compounds, and other key atmospheric species (Crutzen, 1979). NO_x is also directly toxic in high concentrations, plays a significant role in tropospheric ozone production, and serves as a source of NO₃⁻, a key nutrient for ecosystems and significant component of acid rain. NO_x is primarily emitted as nitric oxide (NO) through fossil fuel combustion, biomass burning, lightning and microbial activity in



soils (Seinfeld and Pandis, 2006). NO is rapidly oxidized to nitrogen dioxide (NO₂) through reactions with ozone and peroxy radicals, and NO₂ subsequently photolyzes to reform NO. The interconversion of NO and NO₂ reaches steady-state within a few minutes during the daytime (Crutzen, 1979). The effects of NO_x on urban chemistry, where anthropogenic emissions dominate the NO_x source, have been extensively studied. However, the processes affecting NO_x in forested and agricultural regions are less well-understood.

In forests and agricultural lands, the major source of NO_x is NO emitted as a by-product of microbial denitrification and nitrification (Mckenney et al., 1982; Caranto and Lancaster, 2017). Deposition of NO₂ to plant canopies is thought to be a significant sink of NO_x in forests, substantially reducing the contribution of soil-emitted NO_x to the atmospheric NO_x budget. Jacob and Wofsy (1990) observed low NO_x above the canopy over the Amazon forest during the wet season. Using a 1-D chemical and transport model constrained by observed NO_x and ozone, they concluded that a substantial fraction of soil-NO_x must be absorbed by the canopy. Extrapolation of these ideas to forests with different leaf area indices suggest that 20–50% of the global fraction of soil-emitted NO_x is lost to vegetation (Yienger and Levy, 1995; Lerdau et al., 2000). Using the framework of Jacob and Wofsy (1990), and Yienger and Levy (1995), global atmospheric models have been tuned to describe observed atmospheric NO_x concentrations and tropospheric ozone production using a canopy reduction factor (CRF), an adjustable parameter which accounts for the difference between soil NO emissions and the amount of NO_x ventilated through the canopy (Yienger and Levy, 1995; Ganzeveld et al., 2002a; Vinken et al., 2014). However, CRFs are implemented in an unphysical manner where they act only on soil NO_x emissions and not on other NO_x present in the plant canopy. An improved understanding is needed of the physical processes governing the foliar uptake of NO_x at the ecosystem and leaf scales.

Many studies have also directly observed the leaf-level uptake of NO₂ (Neubert et al., 1993; Rondon and Granat, 1994; Hereid and Monson, 2001; Sparks et al., 2001; Teklemariam and Sparks, 2006; Pape et al., 2009; Chaparro-Suarez et al., 2011; Breuninger et al., 2013). Experiments investigating the mechanism of NO₂ uptake using the ¹⁵N isotope as a tracer have demonstrated that atmospheric NO₂ can be absorbed through the stomata of plant leaves, converted to nitrate (NO₃⁻) and nitrite (NO₂⁻), and eventually assimilated into amino acids (Rogers et al., 1979; Okano and Totsuka, 1986; Nussbaum et al., 1993; Weber et al., 1995; Yoneyame et al., 2003). The mechanism of NO₂ assimilation is diffusion into the stomata followed by dissolution into the aqueous phase and disproportionation to NO₃⁻ and NO₂⁻ in the apoplasm (Lee and Schwartz, 1981a, b). NO₂ can also be transformed to nitrate and nitrite through scavenging by antioxidants, most notably ascorbate (Ramge et al., 1993). The influence of ascorbate on foliar uptake was theoretically calculated by Ramge et al. (1993), and experimentally demonstrated by Teklemariam and Sparks. (2006). The enzyme nitrate reductase converts NO₃⁻ to NO₂⁻ in the cytosol. NO₂⁻ is then transported into the plastids where it is further reduced by the enzyme nitrite reductase to ammonium (NH₄⁺), the product required for amino acid synthesis (Ammann et al., 1995; Tischner, 2000; Teklemariam and Sparks, 2006). Alternatively, NO₂ can deposit directly onto the leaf cuticles or a leaf-surface water film (Burkhardt and Eiden, 1994). However, foliar uptake of NO₂ has been demonstrated to be controlled primarily by the stomata, with deposition to the leaf surface representing only a small fraction of the total NO₂ flux (Thoene et al., 1991; Gessler et al.,



2000; Chaparro-Suarez et al., 2011). Strong correlations have been observed between NO_2 concentrations, stomatal conductances, and the NO_2 deposition flux, suggesting foliar uptake is mainly controlled by stomatal aperture and internal leaf resistances (Johansson, 1987; Thoene et al., 1991; Rondon et al., 1993; Meixner et al., 1997; Chaparro-Suarez et al., 2011; Breuninger et al., 2013).

5 Despite the large body of research that exists on the leaf-level deposition of NO_2 to vegetation, discrepancies still exist of NO_2 exchange rates and the role of mesophilic processes. Many laboratory experiments have failed to measure uptake rates necessary to describe the observed 20–50% reduction of soil-emitted NO_x (Hanson and Lindberg, 1991; Breuninger et al., 2013). Another considerable controversy is the existence of a compensation point—a concentration below which leaves would instead act as a source of NO_x . Compensation points of 0.1–3.2 ppb NO_x have been observed in a
10 number of laboratory chamber studies, suggesting trees instead may serve as a large source of NO_x in forests (Johansson, 1987; Rondon et al., 1993; Hereid and Monson, 2001; Sparks et al., 2001; Teklemariam and Sparks, 2006). Emission of NO at these low NO_x mixing ratios has also been detected in laboratory chamber studies (Wildt et al., 1997; Hereid and Monson, 2001). Observations of NO_x canopy fluxes and atmospheric models conversely predict that trees are substantial sinks of NO_x at concentrations as low as 0.1 ppb, typical of NO_x mixing ratios in remote areas (Jacob and Wofsy, 1990). More recent
15 laboratory studies of leaf level deposition have also questioned the existence of a compensation point (Chaparro-Suarez et al., 2011; Breuninger et al., 2013).

Many laboratory investigations of NO_x foliar exchange have not been performed with instruments sufficiently sensitive or specific to measure uptake of NO_2 at the low NO_x concentrations relevant to forested environments. A commonly used technique for chamber observations of leaf-level NO_2 uptake is the indirect NO_2 measurement technique of
20 chemiluminescence detection of NO (Sparks et al., 2001; Teklemariam and Sparks, 2006; Pape et al., 2009; Chaparro-Suarez et al., 2011; Breuninger et al., 2012; Breuninger et al., 2013). This technique requires photolytic or catalytic conversion of NO_2 to NO , which is either limited by large detection limits greater than 100 ppt (Teklemariam and Sparks, 2006; Pape et al., 2009; Chaparro-Suarez et al., 2011; Breuninger et al., 2012; Breuninger et al., 2013), or may be subject to interferences from higher oxides of nitrogen (Sparks et al., 2001). Further, interferences from the chemiluminescence of alkene + ozone
25 reaction products have also been identified (Reed et al., 2016). These are of particular importance since alkenes make up a substantial fraction of biogenic VOC emissions (e.g. isoprene) (Kesselmeier et al., 2002; Lappalainen et al., 2009; Park et al., 2014; Romer et al., 2016). Even in laboratory settings, where interferences from higher oxides are not of concern, emissions of alkenes from the enclosed leaves may cause substantial interferences. New methods for studying the exchange at the leaf-level, are required to resolve existing discrepancies regarding the foliar uptake rate of NO_2 and the existence of a
30 compensation point.

To understand the leaf-level processes affecting ecosystem scale atmosphere-biosphere NO_x exchange, we have conducted laboratory experiments measuring NO and NO_2 fluxes to the native California tree species *Quercus agrifolia* (Fig. 1) using a branch enclosure system and direct laser-induced fluorescence (LIF) detection of NO_2 (Fig. 2). With the LIF technique, we are able measure NO_x exchange fluxes with high specificity and sensitivity. We investigated the existence of



an NO₂ and NO compensation point and the rate of NO_x foliar uptake under controlled conditions. To our knowledge this is the first leaf-level uptake experiment that has been performed on a North American tree species.

2 Materials and methods

2.1 *Quercus agrifolia* samples

5 NO_x uptake by *Quercus agrifolia* (Coastal Live Oak) was investigated in the laboratory. Three *Quercus agrifolia* individuals were purchased from a local native California plant nursery (Native Here Now Nursery), where the plants were grown from seeds and cuttings collected in Contra Costa County. The tree specimens were grown in a nutrient rich commercial soil mixture (a mixture of Orchard Potting Soil and EB Stone Cactus Mix) at the Jane Grey Research Greenhouse at the University of California, Berkeley. The trees were 2–3 years old when measurements were taken.

10 2.2 Laser-induced fluorescence detection

NO₂ was measured using Laser-Induced Fluorescence (LIF). A blue diode laser (Z-Laser ZM18H3,) centered at a wavelength of 405 nm was focused into each detection cell and made 20 passes in White multipass optical configuration (Fig. 2b)(Thornton et al., 2000; Fuchs et al., 2009). Upon absorption of a visible photon, NO₂ undergoes a transition from the ²A₁ ground to the ²B₂ excited electronic state. The excited NO₂ molecule, is either quenched by collision or emits a red-shifted photon as it relaxes back to ground state (e.g. Thornton et al., 2000). These emitted photons were detected using a red-sensitive photomultiplier tube (PMT) (Hamamatsu H7421-50). To minimize collisional quenching, each detection cell was maintained at a pressure of around 3 torr. Excitation at 405 nm was chosen because it is near the region of maximum absorption in the NO₂ spectrum, and is not subject to interferences from absorption by water vapor or O₃ (Matsumoto and Kajii, 2003).

20 Calibrations were performed every hour by diluting NO (4.97 ppm ± 5%, Praxair) and NO₂ standard gases (5.08 ppm ± 5%, Praxair) to 1–10 ppb in humidified (RH ~60%) zero air. The limit of detection (LOD) for the detection cells is described as:

$$LOD = \frac{S/N}{m} \sqrt{\frac{2b}{t}} \quad (1)$$

where *m* is the slope of the calibration curve constructed from standard dilutions, *b* is the PMT signal at 0 ppb NO or NO₂,
25 *S/N* is the desired signal to noise ratio, and *t* is the time of signal averaging. At a *S/N* of 2 and signal averaging over 5 min, the LOD for detection cells 1–4 was 15 ppt, 4 ppt, 10 ppt, and 30 ppt, respectively. NO₂ in the incoming and outgoing airstreams was measured simultaneously in the first two detection cells. In the second two detection cells, NO was quantitatively converted to NO₂ in the presence of excess ozone, allowing for detection of total NO_x (Fig. 2a). Ozone was produced using an ozone generator (Jelight 600) and flow rates of ozone delivered were adjusted to achieve unity conversion
30 of NO to NO₂.



Much of the previous work on leaf-level NO₂ uptake has been performed using indirect detection of NO₂, in which photolytic or catalytic conversion of NO₂ is followed by chemiluminescence measurement of NO (Sparks et al., 2001; Teklemariam and Sparks, 2006; Pape et al., 2009; Chaparro-Suarez et al., 2011; Breuninger et al., 2012; Breuninger et al., 2013). However, these techniques are either limited by much larger detection limits in the 100–500 ppt range, or are non-specific in their conversion of NO₂ to NO. Our measurements demonstrate a much higher degree of certainty, due largely to a lower detection limit than comparable experiments with specific photolytic conversion and chemiluminescence measurement of NO₂ (Chaparro-Suarez et al., 2011; Breuninger et al., 2012; Breuninger et al., 2013). With the LIF detection of NO₂, we are able to sensitively measure exchange fluxes at low NO_x mixing ratios relevant to forested environments.

2.3 Dynamic chamber system

The NO_x flux measurements were performed with a dynamic branch enclosure system, consisting of a thin transparent double-walled Teflon film (FEP) bag (American Duraflim), which transmits 90% of photosynthetically activated radiation. The chamber was illuminated by an LED diode array of 430–475 nm and 620–670 nm lights (Apollo Horticulture). This light source was selected because it does not emit wavelengths below 420 nm, where NO₂ dissociates, preventing loss of NO₂ to photodissociation and resultant photochemistry. In order to ensure turbulent mixing and minimal aerodynamic and boundary layer resistances, a Teflon-coated fan was installed inside the inner chamber (Meixner et al., 1997; Pape et al., 2009; Breuninger et al., 2013).

Before experiments with *Quercus agrifolia* individuals, the deposition to an empty chamber was measured and background subtracted from subsequent branch measurements. The measured loss of NO₂ to chamber walls was 5% of the NO₂ mixing ratio flowing into chamber. This corresponded to a maximum loss of 0.4 ppb at 8 ppb NO₂. Emission of less than 0.05 ppb NO₂ from the Teflon walls was also observed when chamber lights were turned on with 0 ppb NO₂ flowing through the system. It is likely that the chamber walls buffer uptake of NO₂, but this is a minor effect, as the wall emission observed was a tiny fraction of the measured fluxes.

During measurements, the enclosed branch was exposed to known amounts of either NO₂ or NO mixed with zero air. The inner chamber had an inner diameter of 20 cm, a length of 40 cm, and a total volume of 13 L (American Duraflim 200A Teflon FEP). Flow rates into the inner chamber (Q) during experiments were typically 5 L min⁻¹, creating residence time in the chamber of 3 min. The outer chamber had an inner diameter of 30 cm and a length of 55 cm (American Duraflim 500C20 Teflon FEP). Zero air at a flow rate of 3 L min⁻¹ constantly fumigated the outer bag, serving as a buffer region to ensure the laboratory air, with high mixing ratios of NO_x, did not diffuse into the bag enclosing the branch.

The photosynthetic photon flux density (PPFD) was monitored outside the chamber with a LiCor quantum sensor (LiCor LI-190SA). The flux density measured above the chamber was 1190 μmol m⁻² s⁻¹, approximately the PPF for Berkeley, California at noon during the month of October. This is well above the photon flux required to achieve maximal stomatal aperture for broadleaf evergreen trees (von Caemmerer and Farquhar, 1981; Chaparro-Suarez et al., 2011; Breuninger et al., 2013). We confirmed this assumption by covering the lights with a filter to reduce the intensity by 40%



and monitoring CO₂ and H₂O exchange. No reduction in the exchange rates of these gases were observed. The relative humidity of air entering the chamber was maintained at 50–65% in all experiments by flowing zero air through a bubbler before mixing with NO_x. Measurements of NO_x exchange fluxes occurred under a light/dark cycle with a photoperiod of 12 hours and a temperature of 26/22 ± 2 °C. No change in NO_x uptake was observed when heating the chamber with the lights off or cooling the chamber with the lights on. We therefore expect no significant temperature effects caused by the 4°C difference in temperature between light and dark periods. We also observed a relative humidity increase in the delivered air of about 2% with the lights off, but do not expect this increase to produce any significant changes in NO_x deposition or plant physiology (von Caemmerer and Farquhar, 1981; Chaparro-Suarez et al., 2011).

Exchange of CO₂ and H₂O with the leaves were monitored with a LiCor-6262 H₂O/CO₂ analyzer operating in differential mode. Flows of 0.1 L min⁻¹ of air entering and exiting the chamber were diverted to the LiCor analyzer to measure the CO₂ assimilation and transpiration rates. To measure the CO₂ content and relative humidity of air delivered to the chamber, 0.5 L min⁻¹ of the humidified zero air/NO_x mixture was diverted to a second external 1.5 L cuvette. The temperature and relative humidity of air entering the chamber were measured with a temperature and relative humidity module in the external cuvette (TE Connectivity HTM2500LF). The CO₂ mixing ratios in the external chamber were monitored with a Vaisala CarboCap GMP343 sensor.

2.4 NO_x flux densities

The leaf-level exchange flux of NO or NO₂ (F_{NO_x}) was calculated according to Eq. 2:

$$F_{NO_x} = \frac{Q \cdot (C_o - C_i)}{A} \quad (2)$$

where Q is the flow rate (m³ s⁻¹), A is the enclosed leaf area (m²), C_o is the concentration leaving the chamber, and C_i is the concentration entering the chamber (nmol m⁻³). The calculated flux is related to a deposition velocity (Vd_{NO_x}) by Eq. 3:

$$F_{NO_x} = -Vd_{NO_x} \cdot (C_o + C_{comp}) \quad (3)$$

where C_{comp} is the compensation point, the concentration of NO₂ below which the tree would instead act as a source of NO_x.

The deposition velocities were calculated through weighted least squares regression of calculated fluxes and outlet NO_x concentrations (C_o). The absolute value of the slope of the regression line was equal to the deposition velocity, with the x-intercept representing the compensation point concentration. The precision error in the NO_x exchange flux (σ_F) was calculated through propagation of the error in the inlet (σ_{C_i}) and outlet (σ_{C_o}) concentrations (Eq. 4).

$$\sigma_F = \frac{Q}{A} \sqrt{\sigma_{C_i}^2 + \sigma_{C_o}^2} \quad (4)$$

σ_{C_i} and σ_{C_o} were estimated as the larger of the error in the calibration slopes and the standard deviation of the 5 min signal average. From observations in daily deviations of the flow rate and error in measured leaf area using the ImageJ software (Schneider et al., 2012), we estimate the error in $\frac{Q}{A}$ to be a maximum of 0.005 cm s⁻¹. This usually was only a minor contribution to the total error in the NO_x exchange flux.



The calculated deposition velocity was used to find the total resistance to deposition, R , via Eq. 5.

$$Vd_{NO_x} = \frac{1}{R} \quad (5)$$

The total resistance is described by the canopy stomatal resistance model (Baldocchi et al., 1987) and defined in Eq. 6–7.

$$R = R_a + R_b + R_{leaf} \quad (6)$$

$$5 \quad R_{leaf} = \left(\frac{1}{R_{cut}} + \frac{1}{R_{st} + R_m} \right)^{-1} \quad (7)$$

where R_{leaf} is the total leaf resistance and R_a , R_b , R_{cut} , R_{st} , and R_m are the aerodynamic, boundary layer, cuticular, stomatal, and mesophilic resistances, respectively. The aerodynamic resistance is characterized by the micrometeorology above a surface and is dependent upon the wind speed and turbulence of air flow. The boundary layer resistance describes the diffusion of a molecule through a shallow boundary of air above a surface and is dependent on microscopic surface properties, diffusivity of the gas species, wind speed, and turbulence of air flow (Baldocchi et al., 1987). R_{cut} , R_{st} , and R_m are the resistances associated with deposition to the leaf cuticles or through the stomata, and are dependent upon plant physiology.

The chamber fan, installed to create turbulent mixing, allowed for the assumption that R_a was negligible (Pape et al., 2009; Breuninger et al., 2012). R_b is chamber-specific, and has typically not been measured in previous chamber experiments of NO_2 leaf-level deposition (Chaparro-Suarez et al., 2011; Breuninger et al., 2012; Breuninger et al., 2013). R_b was experimentally measured in this study by placing a tray of activated carbon into the chamber (assumed to have zero surface resistance to deposition of NO_2), and calculating the deposition flux of NO_2 . The leaf components to the total deposition resistance were determined through dark and light experiments. During dark experiments, the stomata were closed (confirmed with measurements of CO_2 and H_2O exchange), and the deposition observed was assumed to be entirely driven by deposition to the cuticles.

3 Results

3.1 Determination of the boundary resistance R_b

To estimate the chamber boundary layer resistance and test the assumption that $R_b \ll R_{leaf}$, a dish of activated carbon, which theoretically has zero chemical resistance to deposition of NO_2 , was placed inside the chamber. The boundary layer resistance was considered to be the only component of the total resistance to deposition. The deposition velocity of NO_2 to activated carbon was measured as $0.52 \pm 0.06 \text{ cm s}^{-1}$, corresponding to a boundary layer resistance to NO_2 deposition of $1.94 \pm 0.02 \text{ s cm}^{-1}$ (Fig. 3). This boundary resistance is approximately double what was measured by Pape et al. 2009—a reasonable difference given differences in chamber design (Fig. 2). The R_b for NO_2 was scaled with the ratio of diffusivities of NO_2 and NO in air to obtain the resistance to deposition of NO of $2.59 \pm 0.03 \text{ s cm}^{-1}$.



The boundary resistance was also estimated in an additional experiment (not shown) in which a de-ionized water-soaked Whatman No. 1 filter paper was placed inside the chamber and the evaporation of water vapor into the chamber filled with dry zero air was measured. The emission flux of water vapor from the filter paper was calculated in a similar manner to that of NO_x deposition flux (Eq. 2). The conductance to water vapor was then calculated via:

$$5 \quad \frac{Q \cdot (P_{H_2O})}{A} = g_w (P_{sat} - P_{H_2O}) \quad (8)$$

where P_{H_2O} is the partial pressure of water vapor inside the chamber, P_{sat} is the saturation vapor pressure at the temperature in the chamber, and g_w is the conductance to water vapor. The measured conductance to water vapor was scaled with the ratio of diffusivities of NO₂ to water vapor (D_{NO_2}/D_{H_2O}) and inverted to find the NO₂ boundary layer resistance:

$$R_b = \frac{D_{H_2O}}{D_{NO_2}} \frac{1}{g_w} \quad (9)$$

- 10 The boundary resistance to NO₂ deposition by this method was found to be 2 s cm⁻¹, essentially identical to the measurement on the activated-carbon.

3.2 NO_x deposition velocity and compensation point concentration

The deposition velocities and compensation points were respectively calculated as the slope x-axis intercept of the regression line between NO_x exchange flux and chamber NO_x concentrations (Fig. 4). The detection limit was a dominant source of error in the estimation of the NO exchange flux and compensation point. The large relative uncertainties in NO flux measurements were caused by the much slower deposition of NO compared with that of NO₂, inhibiting our ability to observe the very small changes between the NO concentration in the chamber and the incoming airstream (Fig. 4). Additional uncertainty in NO₂ flux measurements because of enhanced quenching of NO₂ by water vapor should be minimal, as calibrations and measurements were performed at equivalent relative humidities. However, transpiration of the enclosed leaves caused the absolute humidity within chamber to be enhanced by 0.3–0.5% relative to the incoming airstream. We expect this to result in a maximum error in calculated NO₂ mixing ratios of 1–1.75% (Thornton et al., 2000), resulting in maximum errors in the calculated fluxes and deposition velocities of 2% and 4%, respectively. This 4% error in the calculated deposition velocity during lights-on experiments is considerably less than the uncertainty of the linear fit (Fig. 4).

Correlation coefficients, deposition velocities, compensation points, and statistical testing of the compensation point for NO₂ and NO deposition are shown in Table 1 and Table 2, respectively, and were calculated according to Breuninger et al. (2013). For NO₂ experiments, only one dark and one light experiment with *Quercus agrifolia 1*, was found to have a statistically significant ($\alpha = 0.05$) non-zero intersection with the x-axis (Table 1). The range of C_{comp} measured were -0.02–0.300 ppb NO₂, with probabilities of $C_{comp} = 0$ ranging from 10.3–91.6% (excluding the two *Quercus agrifolia 1* experiments) (Table 1). Conversely, all three *Quercus agrifolia* individuals during all dark and light NO deposition experiments demonstrated compensation points significantly above zero, ranging from 0.74–3.8 ppb NO. For all light and



dark experiments the average compensation point for was calculated as 0.84 ± 0.32 ppb NO and 2.4 ± 1.1 ppb NO, respectively (Table 2).

Student's *t* tests, (not shown), demonstrated that deposition velocities and compensation points measured during NO and NO₂ lights on and off experiments were not significantly different (to the $\alpha=0.05$ confidence level) between different *Quercus agrifolia* individuals. Deposition velocities for NO₂ light experiments were between 0.08 and 0.18 cm s⁻¹, with a deposition of 0.123 ± 0.007 cm s⁻¹ calculated from the regression of all light experiments. Dark experiments resulted in deposition velocities between 0.013 and 0.022 cm s⁻¹, with a deposition velocity of 0.015 ± 0.001 cm s⁻¹ calculated from the regression of all dark experiments (Table 1). NO demonstrated much slower deposition, with deposition velocities from all light and dark experiments calculated as 0.012 ± 0.002 cm s⁻¹ and 0.005 ± 0.002 cm s⁻¹, respectively (Table 2). Despite the large compensation point measured for NO, the leaf emission fluxes of NO were a maximum of only 0.8 pmol m⁻² s⁻¹ at 0.1 ppb NO, approximately half of the deposition flux measured for NO₂ at 0.1 ppb, making *Quercus agrifolia* a large net sink of NO_x.

3.3 Resistances to Leaf-level NO_x deposition.

The deposition velocity measured from linear regression of NO_x exchange fluxes and NO_x chamber concentrations is the inverse of the total resistance to deposition (Eq. 6), with R_a assumed to be zero. The total resistance in the chamber is thus:

$$R = R_b + \left(\frac{1}{R_{cut}} + \frac{1}{R_{st} + R_m} \right)^{-1} \quad (10)$$

The leaf resistance to deposition can then be found by subtracting the boundary layer resistance from the total resistance. Total leaf resistances, R_{leaf} , were calculated using the boundary layer resistances for NO₂ and NO of 1.94 ± 0.02 s cm⁻¹ and 2.59 ± 0.03 s cm⁻¹, respectively. During the dark experiments, R_{leaf} is equal to R_{cut} , and the deposition velocity measured was estimated as the inverse of the sum of the boundary and cuticular resistance. After calculation of R_{cut} from dark experiments, the sum of the stomatal and mesophilic contributions to the total leaf resistance was determined. The boundary, cuticular, and summed stomatal and mesophilic resistances are shown in Table 3.

It must be noted that it is possible that the stomata were not entirely closed during dark experiments. Evidence exists that nocturnal stomatal conductance can be large enough to allow for significant transpiration (Dawson et al., 2007), and small (within the range of uncertainty observed for the LICOR-6262) emission of water vapor during dark experiments was measured. However, even if all the deposition during dark experiments was stomatal, this would cause only a 0.5 s cm⁻¹ reduction in the calculated R_{st} for NO₂, less than the uncertainty from the error in the measured deposition velocity (~10% error). The cuticular resistances reported here during dark experiment are nonetheless atmospherically relevant to nighttime NO_x deposition.



4 Discussion

4.1 NO_x deposition velocities and compensation points

The strong linear dependence between NO₂ fluxes and NO₂ chamber concentrations that we observe is consistent with previous observations that NO₂ exchange is largely driven by NO₂ concentration differences between the atmosphere and gaseous phase of the leaf (Rondon and Granat, 1994; Gessler et al., 2000; Hereid and Monson, 2001; Sparks et al., 2001; Teklemariam and Sparks, 2006; Pape et al., 2009; Chaparro-Suarez et al., 2011; Breuninger et al., 2012). Our measurements of NO₂ stomatal resistance parameters for *Quercus agrifolia* represents a stomatal deposition velocity (inverse of $R_{st} + R_m$) of $0.14 \pm 0.02 \text{ cm s}^{-1}$. This value is similar to the range of 0.1–0.15 cm s^{-1} that Chaparro-Suarez et al. (2011) found for two European oak tree species, *Quercus robur* and *Quercus ilex*. The deposition velocity measured here for *Quercus agrifolia* is also much larger than 0.007–0.042 cm s^{-1} range found for Norway spruce (*Picea abies*) by Breuninger et al. (2012), but surprisingly comparable, given the differences in plant species, to the 0.12 cm s^{-1} deposition velocity found for maize (*Zea mays*) by Hereid and Monson (2001). We also find here a NO₂ flux at 5 ppb of 0.2 $\text{nmol m}^{-1} \text{ s}^{-1}$, similar in magnitude to the 0.1 $\text{nmol m}^{-1} \text{ s}^{-1}$, 0.15–1.5 $\text{nmol m}^{-1} \text{ s}^{-1}$, and 0.18 $\text{nmol m}^{-1} \text{ s}^{-1}$ fluxes measured for *Fagus sylvatica* (Gessler et al., 2000), tropical Panamanian native trees (Sparks et al., 2001), and periwinkle (*Catharanthus roseus*) (Teklemariam and Sparks, 2006), respectively.

Resistance parameters reported above for NO deposition to *Quercus agrifolia* represent a stomatal deposition velocity of $0.007 \pm 0.002 \text{ cm s}^{-1}$ and cuticular deposition velocity of $0.005 \pm 0.001 \text{ cm s}^{-1}$. This observation of very minor NO uptake—at least an order of magnitude less than that of NO₂ uptake—is also consistent with previous observations (Hanson and Lindberg, 1991; Hereid and Monson, 2001; Teklemariam and Sparks, 2006). We also detected a statistically significant NO compensation point, with low emissions up to 8 $\text{pmol m}^{-2} \text{ s}^{-1}$ observed below 1 ppb. These observations are similar to the 8–14 $\text{pmol m}^{-2} \text{ s}^{-1}$ emission fluxes of NO reported by Hereid and Monson (2001) and Teklemariam and Sparks (2006) at low NO_x concentrations.

No significant NO₂ compensation point was found for our measurements of *Quercus agrifolia* NO_x uptake. Many previous studies have reported NO₂ compensation points, ranging from 0.1–3.0 ppb, implicating trees as a constant source of NO_x in forest ecosystems (Gessler et al., 2000; Hereid and Monson, 2001; Sparks et al., 2001; Teklemariam and Sparks, 2006). Our findings of a lack of NO₂ compensation point support field observations and modeling studies that have recognized NO₂ dry deposition to vegetation as an important NO_x loss process in forests (Jacob and Wofsy, 1990; Ganzeveld et al., 2002b; Geddes and Murphy, 2014). Our results also support the works of Chaparro-Suarez et al. (2011) and Breuninger et al. (2013), who did not find evidence of an NO₂ compensation point.

The primary difference in our experimental setup, compared to previous dynamic chamber studies that have found a NO₂ compensation point, is the use of a direct NO₂ measurement technique. Measurements of a significant NO₂ compensation point have mostly been obtained using techniques requiring conversion of NO₂, followed by chemiluminescence detection of NO (Gessler et al., 2000; Hereid and Monson, 2001; Sparks et al., 2001; Teklemariam and



Sparks, 2006). Such methods have utilized either non-specific photolytic (Gessler et al., 2000; Hereid and Monson, 2001), luminol (Sparks et al., 2001), or catalytic conversion (Teklemariam and Sparks, 2006) techniques, which may have also resulted in the conversion of PAN, HONO, HNO₃, and other organic nitrates, as well as interferences from alkene + ozone reactions. If any of these interfering compounds are not excluded from the chamber system, or form from reactions of biogenic emissions, this would cause an enhancement in observed NO₂ compensation point, and a suppression of observed deposition velocity. Additionally, previous chamber measurements have sometimes employed chamber setups that would let in a substantial amount of UV light, yet did not exclude photochemical reactions between NO₂, NO, and O₃. Such corrections are excluded here because of our use of chamber lights with only wavelengths above 420 nm. To avoid this issue, other experiments have instead involved a setup including a simultaneously measured blank chamber, which would theoretically allow for correction for any reactions resulting from photolysis of NO₂, O₂, or O₃ (Gessler et al., 2000; Hereid and Monson, 2001). Such corrections might be complicated by secondary chemistry not present in our experiments.

4.2 Implication for canopy NO_x loss

Resistance parameters reported above (Table 3) were used in a 1-D seven-layer multibox model representing chemical reactions, vertical transport, and leaf-level processes scaled to the canopy level to assess the impacts of NO_x deposition velocities on the NO_x lifetime and fluxes. The model is constructed in a manner similar to Wolfe and Thornton (2011). Details will be presented elsewhere. The 1-D model was run for meteorological conditions representing the native habitat of *Quercus agrifolia* and two different leaf area indices (LAI) approximately representing the lower and upper limits of LAI found in California oak woodlands. As shown in Fig. 5a and 5b, the model predicts NO_x deposition to *Quercus agrifolia* accounts for 3%–7% of the total NO_x loss within the boundary layer if the only source of NO_x is emissions from the soil. This represents a total NO_x lifetime of 7–7.5 hours in the boundary layer, and a lifetime to deposition of 4–11 days in the boundary layer and 0.5–1.2 hours below the canopy. Under these scenarios approximately 15–30% of soil-emitted NO_x is removed in the canopy (Fig. 6)—on the lower end of the range of 25–80% reduction observed in field studies (Jacob and Wofsy, 1990; Lerdaun et al., 2000; Ganzeveld et al., 2002a; Min et al., 2014).

The coastal regions of California where *Q. agrifolia* is found frequently experience much higher NO_x mixing ratios of 10–50 ppb. This is particularly important for oak woodlands of the San Francisco Bay and near Los Angeles areas, where anthropogenic emissions from nearby urban centers are the majority of the NO_x source. To account for this extra NO_x source, additional model runs were done with an added term accounting for NO_x advection from a more concentrated upwind source ($C_{NO_x(adv)}$), with advection treated as a simple mixing process:

$$\left(\frac{dC_{NO_x}}{dt}\right) = -k_{mix}(C_{NO_x} - C_{NO_x(adv)}) \quad (11)$$

where $k_{mix} = 0.3 \text{ h}^{-1}$ and $C_{NO_x(adv)}$ is 10 ppb.



In this case, deposition to *Q. agrifolia* could account for 10–22% of the total NO_x loss (Fig. 5c,d), representing a lifetime to deposition of 5–14 days in the boundary layer and a total NO_x lifetime of 28–33 hours. Deposition in this higher NO_x scenario decreased the total NO_x lifetime by 3–8 hours, compared with a no-deposition case.

5 Conclusions

5 This work constitutes the first measurements of NO₂ and NO foliar deposition resistance parameters for a North American tree species. We report observations of leaf-level resistances to NO₂ and NO deposition, corresponding to total deposition velocities of NO₂ and NO of 0.123 ± 0.007 cm s⁻¹ and 0.012 ± 0.002 cm s⁻¹ in the light and 0.015 ± 0.001 cm s⁻¹ and 0.005 ± 0.002 cm s⁻¹ in the dark, respectively. No compensation point was observed for NO₂, but compensation points of 0.74–3.8 ppb were recorded for NO. The magnitude of NO emission below the compensation point was significantly less than the magnitude of NO₂ uptake in the same concentration range, making *Q. agrifolia* an overall large net sink of NO_x. The observed deposition is large enough to explain canopy reduction factors observed in canopy-level studies, but is at the lower end of estimated global CRFs. The results of the 1-D multibox model demonstrate that the deposition observed accounts for 5–20% of NO_x removal with a NO_x lifetime to deposition of 0.5–1.2 hours beneath the canopy of a California oak woodland. We show that foliar deposition of NO_x represents a significant removable mechanism of NO_x and can have a large impact on NO_x mixing ratios and fluxes in such ecosystems. Further investigations of NO₂ deposition to a larger variety of plant species under a range of environmental conditions are needed to accurately understand the global impacts of NO₂ deposition across diverse ecosystems.

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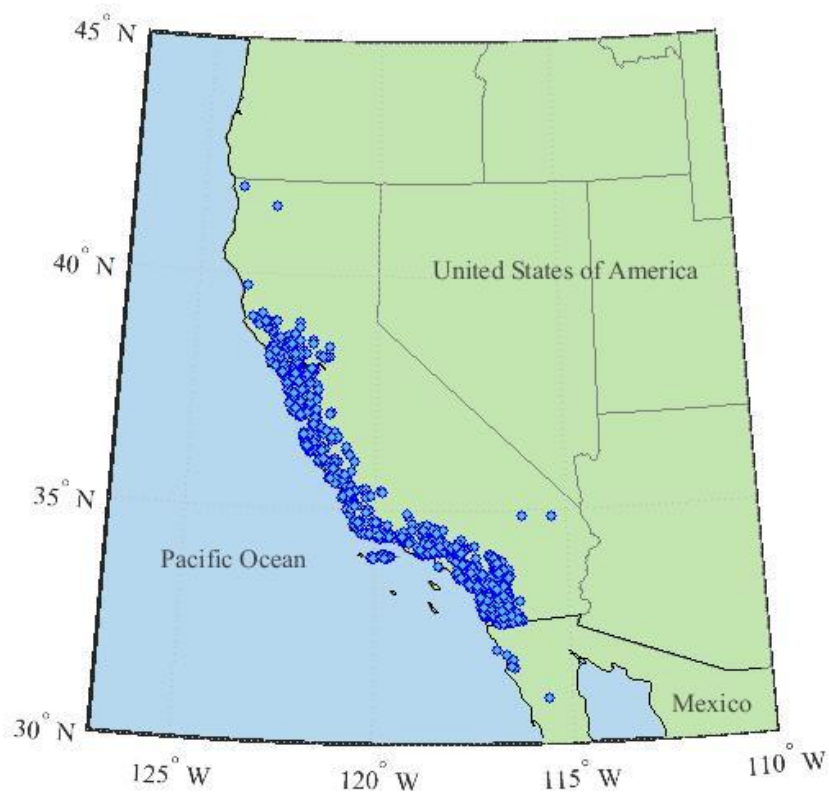


Figure 1: Species distribution map of *Quercus agrifolia*. Each dot represents an observation of *Q. agrifolia* occurrence. Data provided by the participants of the Consortium of California Herbaria.

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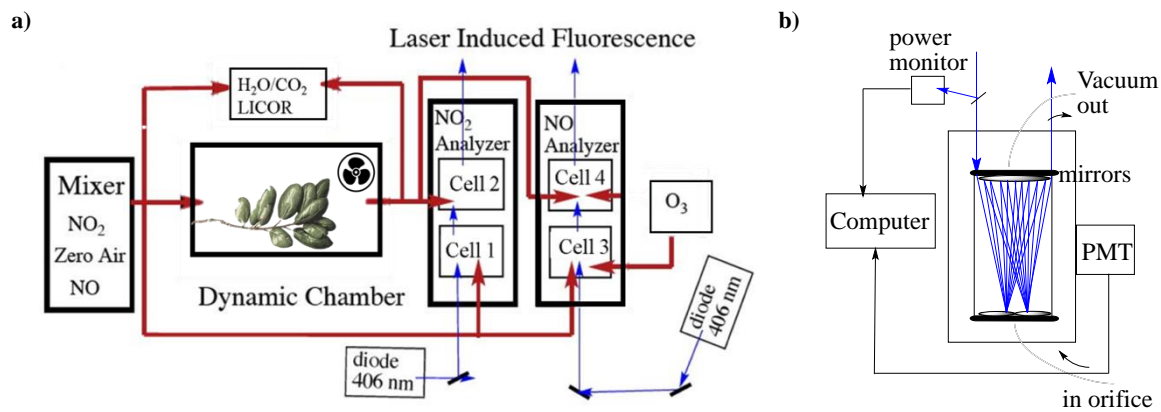


Figure 2: Schematic of the experimental dynamic chamber (a) and laser-induced fluorescence detection (b) setups.

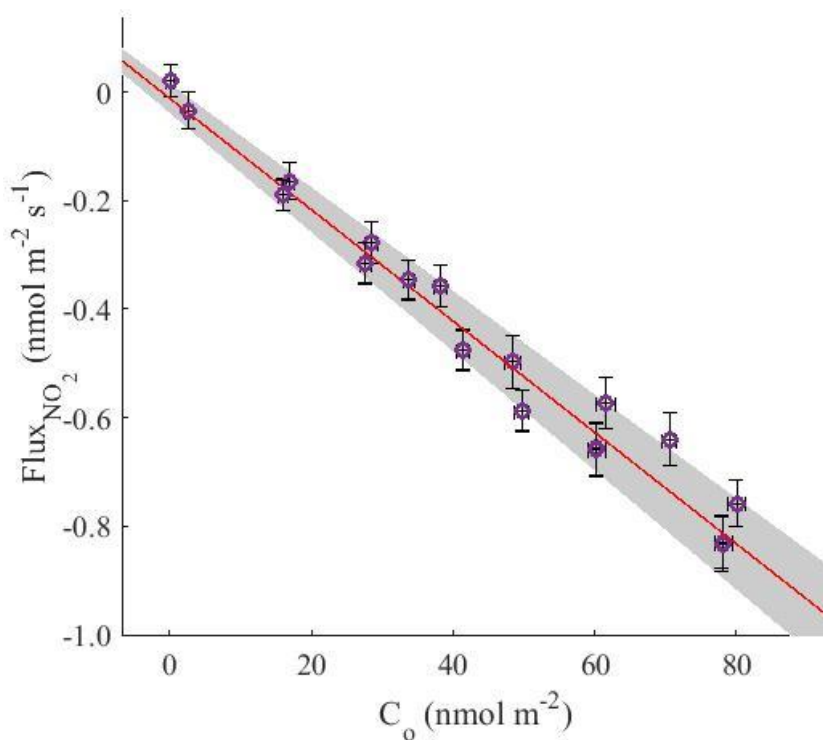


Figure 3: Flux to a 5.1 cm diameter dish filled with activated charcoal. The chemical surface resistance to deposition is approximately zero, so the deposition velocity for deposition of NO_2 to the surface of the charcoal dish is the reciprocal of the boundary layer resistance. The line of best fit is $(0.51 \pm 0.032)C_o$, where C_o is the concentration of NO_2 in the outgoing airstream.

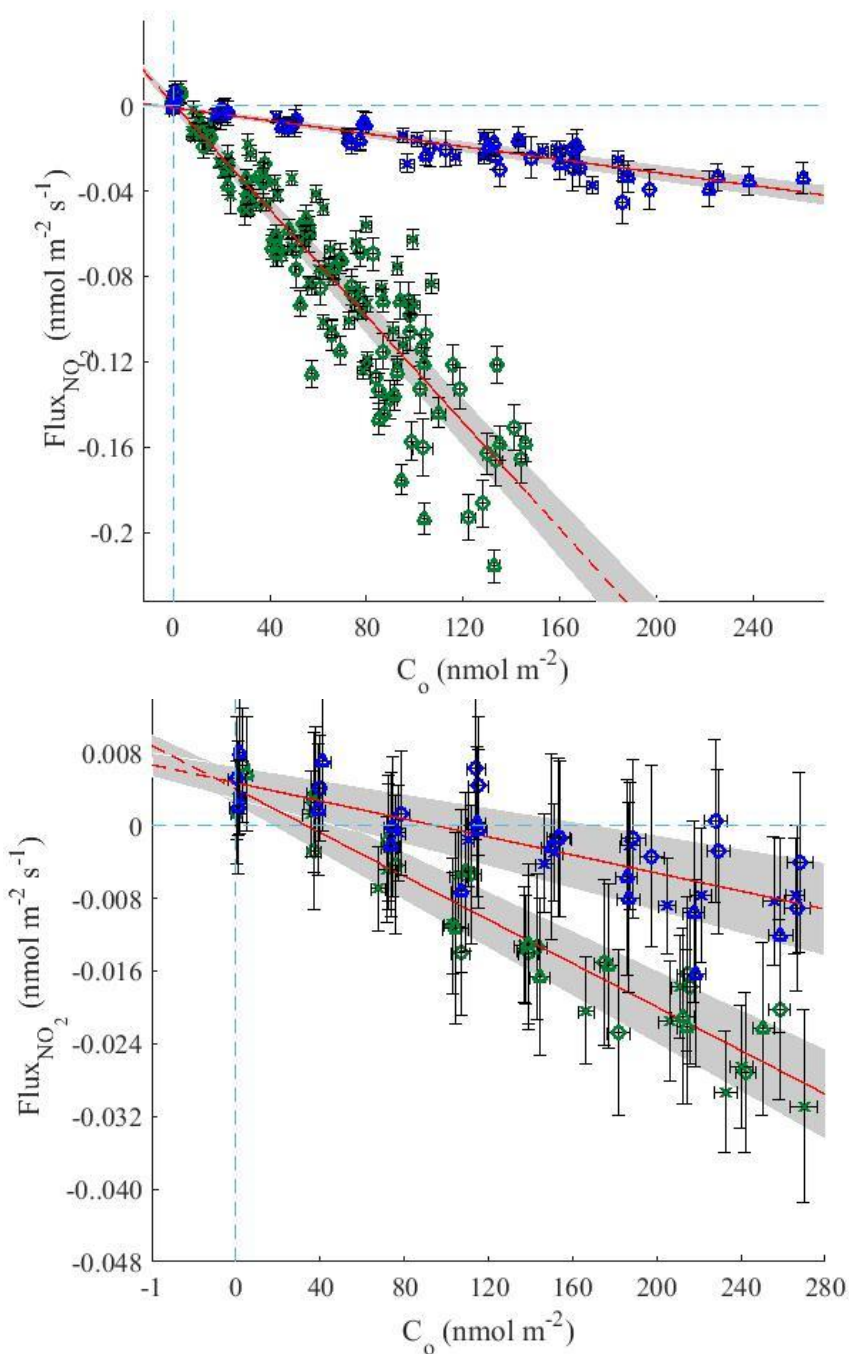


Figure 4: NO₂ (top) and NO (bottom) fluxes versus the outlet concentrations for all *Quercus agrifolia* individuals with the chamber lights on (green) and off (blue). The line of best fit is shown in red and was calculated to minimize the weighted residuals in both the x- and y- axis. The blue dotted line shows where flux is zero. A significantly positive ($\alpha = 0.5$) x-intercept occurs for NO, but not NO₂ experiments.

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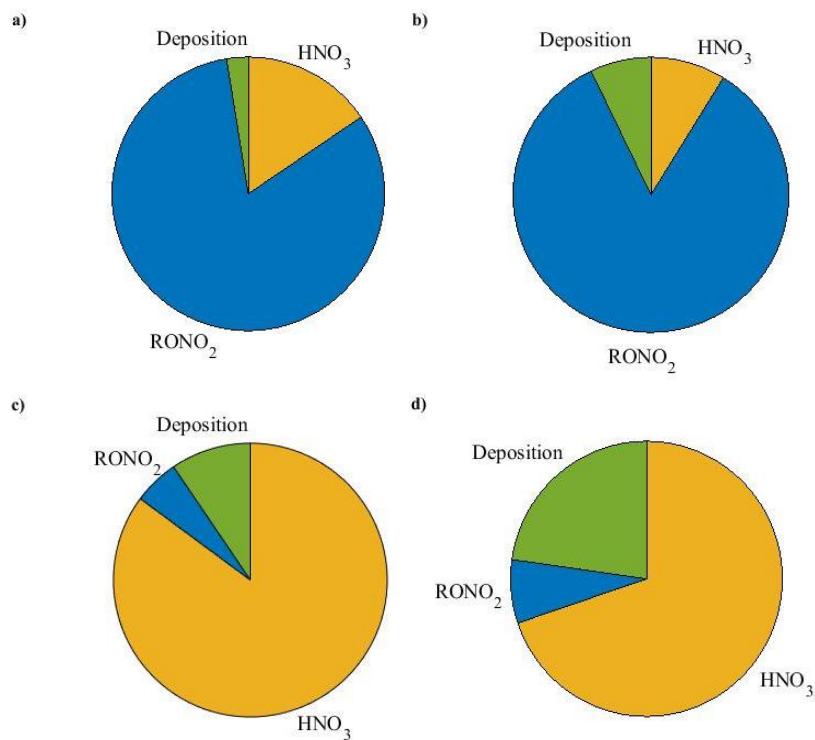
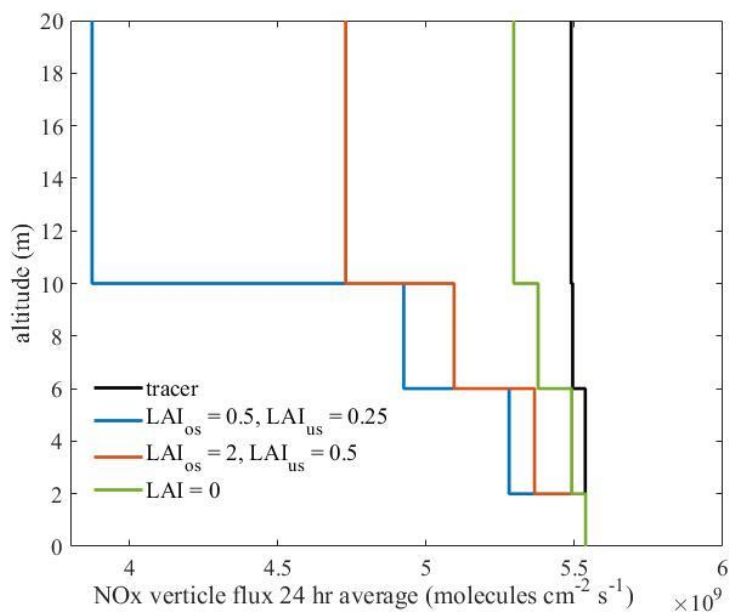


Figure 5: Model predictions of the fraction of NO_x loss to alkyl nitrate formation, nitric acid formation, and deposition in a *Q. agrifolia* woodland. The model was run using scenarios with only soil emissions and LAI of 1 m²/m² (a), only soil emissions and LAI of 3 m²/m² (b), C_{NO_x(adv)} = 10 ppb and LAI of 1 m²/m² (c), and C_{NO_x(adv)} = 10 ppb and LAI of 3 m²/m² (d).

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5 Figure 6: 24 hr average vertical fluxes of NO_x predicted by the 1-D multibox model for a California oak woodland using the leaf resistances measured in this study. Model runs were conducted for a low (red) and high (blue) LAI cases and for a no deposition scenario (green).

**Table 1: Parameters of NO₂ bi-variate linear least-square fitting regression analysis**

run	N	R ²	[NO ₂] _{comp} (ppb)	P([NO ₂] _{comp} =0) %	V _{dep} cm s ⁻¹
<i>Q.agrifolia</i> 1, light					
1	13	0.979	0.056 ± 0.013	42.7	0.10 ± 0.013
2	13	0.950	0.046 ± 0.19	63.7	0.12 ± 0.023
3	16	0.978	0.099 ± 0.086	3.87	0.15 ± 0.016
4	16	0.958	0.077 ± 0.14	28.7	0.12 ± 0.021
All	58	0.927	0.080 ± 0.10	11.6	0.12 ± 0.012
<i>Q.agrifolia</i> 2, light					
1	16	0.963	0.10 ± 0.12	10.3	0.08 ± 0.011
2	5	0.969	-0.01 ± 0.96	83.8	0.12 ± 0.014
3	9	0.997	0.023 ± 0.032	20.3	0.16 ± 0.011
4	16	0.9736	-0.019 ± 0.074	61.9	0.14 ± 0.017
5	15	0.979	0.015 ± 0.082	72.7	0.12 ± 0.014
All	71	0.845	-0.0077 ± 0.091	91.6	0.11 ± 0.014
<i>Q.agrifolia</i> 3, light					
1	11	0.969	0.016 ± 0.18	87.4	0.12 ± 0.024
2	15	0.961	0.074 ± 0.16	39.1	0.18 ± 0.029
3	5	0.990	0.30 ± 0.20	5.9	0.12 ± 0.038
All	31	0.830	0.019 ± 0.064	77.6	0.14 ± 0.029
All <i>Q.agrifolia</i> , light	160	0.885	0.030 ± 0.072	41.3	0.123 ± 0.0092
<i>Q.agrifolia</i> 1, dark					
1	16	0.964	0.056 ± 0.14	0.9*	0.022 ± 0.0034
<i>Q.agrifolia</i> 2, dark					
1	16	0.858	-0.16 ± 0.47	50.8	0.016 ± 0.0050
2	12	0.932	-0.34 ± 0.40	11.8	0.013 ± 0.0038
All	28	0.853	-0.24 ± 0.32	15.6	0.015 ± 0.0030
<i>Q.agrifolia</i> 3, dark					
1	14	0.900	-0.30 ± 0.48	24.1	0.015 ± 0.0042
2	11	0.909	-0.001 ± 0.69	36.7	0.015 ± 0.0057
All	25	0.898	-0.22 ± 0.38	25.3	0.014 ± 0.0029
All <i>Q.agrifolia</i> , dark	69	0.881	-0.16 ± 0.24	12.2	0.015 ± 0.0018

* Significant non-zero compensation point

**Table 2: Parameters of NO bi-variate linear least-square fitting regression analysis**

run	N	R ²	[NO ₂] _{comp} (ppb)	P([NO ₂] _{comp} =0)	V _{dep}
<i>Q. agrifolia</i> 1					
light	17	0.874	0.74 ± 0.65	3.5*	0.011 ± 0.0032
dark	13	0.699	3.8 ± 2.2	0.52*	0.0040 ± 0.0025
<i>Q. agrifolia</i> 1					
light	14	0.954	0.76 ± 0.49	0.92*	0.013 ± 0.0027
dark	10	0.866	1.7 ± 1.0	1.1*	0.0046 ± 0.0018
<i>Q. agrifolia</i> 1					
light	12	0.936	1.3 ± 0.60	0.17*	0.0123 ± 0.0029
dark	15	0.803	2.0 ± 1.0	2.5*	0.0074 ± 0.0033
All <i>Q. agrifolia</i>					
light	13	0.908	0.84 ± 0.32	<0.01*	0.012 ± 0.0015
dark	13	0.602	2.4 ± 1.1	<0.01*	0.0050 ± 0.0016

*Significant non-zero compensation point



Table 3: Summary of deposition resistance parameters of *Quercus agrifolia*

gas	R_b $s\ cm^{-1}$	R_{cut} $s\ cm^{-1}$	$R_{st} + R_m$ $s\ cm^{-1}$
NO ₂	1.94 ± 0.02	63 ± 8	6.9 ± 0.9
NO	2.59 ± 0.03	200 ± 60	140 ± 40