

Measurements of NO and NO₂ exchange between the atmosphere and *Quercus agrifolia*

Erin R. Delaria¹, Megan Vieira,¹ Julie Cremieux², Ronald. C. Cohen^{1,3}

¹Department of Chemistry, University of California, Berkeley, 94720, USA

5 ²Department of Chemistry, Université Pierre et Marie Curie, Paris, France

³Department of Earth and Planetary Science, University of California, Berkeley, 94720, USA

Correspondence to: R. C. Cohen (rccohen@berkeley.edu)

Abstract. NO₂ foliar deposition through the stomata of leaves has been identified as a significant sink of NO_x within a forest canopy. In this study, we investigated NO₂ and NO exchange between the atmosphere and the leaves of the native California 10 oak tree *Quercus agrifolia* using a branch enclosure system. NO₂ detection was performed with laser-induced fluorescence (LIF), which excludes biases from other reactive nitrogen compounds and has a low detection limit of 5–50 ppt. We performed both light and dark experiments with concentrations between 0.5–10 ppb NO₂ and NO under constant ambient 15 conditions. Deposition velocities for NO₂ during light and dark experiments were $0.123 \pm 0.009 \text{ cm s}^{-1}$ and $0.015 \pm 0.001 \text{ cm s}^{-1}$, respectively. Much slower deposition was seen for NO, with deposition velocities of $0.012 \pm 0.002 \text{ cm s}^{-1}$ and $0.005 \pm 0.002 \text{ cm s}^{-1}$ measured during light and dark experiments, respectively. This corresponded to a summed resistance of the 20 stomata and mesophyll of $6.9 \pm 0.9 \text{ s cm}^{-1}$ for NO₂ and $140 \pm 40 \text{ s cm}^{-1}$ for NO. No significant compensation point was detected for NO₂ uptake, but compensation points ranging from 0.74–3.8 ppb were observed for NO. NO₂ and NO 25 deposition velocities reported here are comparable both with previous leaf-level chamber studies and inferences from canopy-level field measurements. In parallel with these laboratory experiments, we have constructed a detailed 1-D atmospheric model to assess the contribution of leaf-level NO_x deposition to the total NO_x loss and NO_x canopy fluxes. Using the leaf uptake rates measured in the laboratory, these modeling studies suggest loss of NO_x to deposition in a California oak woodland competes with the pathways of HNO₃ and RONO₂ formation, with deposition making up 3–22% of the total NO_x loss. Additionally, foliar uptake of NO_x at these rates could account for ~15–30% canopy reduction of soil NO_x 30 emissions.

25 1 Introduction

Nitrogen oxides (NO_x \equiv NO + NO₂) are a group of highly reactive trace gases that control the oxidative capacity of the atmosphere by regulating the amounts of ozone, hydroxyl radicals, volatile organic compounds, and other key atmospheric species (Crutzen, 1979). NO_x is also directly toxic in high concentrations, plays a major role in tropospheric ozone production, and serves as a source of NO₃⁻, a key nutrient for ecosystems and component of acid rain. NO_x is primarily 30 emitted as nitric oxide (NO) through fossil fuel combustion, biomass burning, lightning and microbial activity in soils

(Seinfeld and Pandis, 2006). NO is rapidly oxidized to nitrogen dioxide (NO_2) through reactions with ozone and peroxy radicals, and in the daytime NO_2 subsequently photolyses to reform NO. The interconversion of NO and NO_2 reaches steady-state within a few minutes during the daytime (Crutzen, 1979). The effects of NO_x on urban chemistry, where anthropogenic emissions dominate the NO_x source, have been extensively studied. However, the processes affecting NO_x in 5 forested and agricultural regions are less well-understood.

In forests and agricultural lands, the major source of NO_x is NO emitted as a by-product of microbial denitrification and nitrification (McKenney et al., 1982; Caranto and Lancaster, 2017). Deposition of NO_2 to plant canopies is thought to be an important sink of NO_x in forests, substantially reducing the contribution of soil-emitted NO_x to the atmospheric NO_x budget. Jacob and Wofsy (1990) observed low NO_x above the canopy over the Amazon forest during the wet season. Using a 10 1-D chemical and transport model constrained by observed NO_x and ozone, they concluded that a substantial fraction of soil- NO_x must be absorbed by the canopy. Extrapolation of these ideas to forests with different leaf area indices suggest that 20–50% of the global fraction of soil-emitted NO_x is lost to vegetation (Yienger and Levy, 1995; Lerdau et al., 2000). Using the framework of Jacob and Wofsy (1990), and Yienger and Levy (1995), global atmospheric models have been tuned to 15 describe observed atmospheric NO_x concentrations and tropospheric ozone production using a canopy reduction factor (CRF). The CRF is an adjustable parameter which accounts for the difference between soil NO emissions and the amount of NO_x ventilated through the canopy (Yienger and Levy, 1995; Vinken et al., 2014). However, CRFs are implemented in an unphysical manner where they act only on soil NO_x emissions and not on other NO_x present in the plant canopy. An improved understanding is needed of the physical and biochemical processes governing the foliar uptake of NO_x at the ecosystem and leaf scales.

20 Many studies have also directly observed the leaf-level uptake of NO_2 (Neubert et al., 1993; Rondon and Granat, 1994; Hereid and Monson, 2001; Sparks et al., 2001; Teklemariam and Sparks, 2006; Pape et al., 2009; Chaparro-Suarez et al., 2011; Breuninger et al., 2013). Isotope labeling experiments investigating the mechanism of NO_2 uptake have demonstrated that atmospheric NO_2 can be absorbed through the stomata of plant leaves, converted to nitrate (NO_3^-) and nitrite (NO_2^-), and eventually assimilated into amino acids (Rogers et al., 1979; Okano and Totsuka, 1986; Nussbaum et al., 25 1993; Weber et al., 1995; Yoneyama et al., 2003). The mechanism of NO_2 assimilation is diffusion into the stomata followed by dissolution into the aqueous phase and disproportionation to NO_3^- and NO_2^- in the apoplasm (Lee and Schwartz, 1981a, b). NO_2 can also be transformed to nitrate and nitrite through scavenging by antioxidants, most notably ascorbate (Ramge et al., 1993). The influence of ascorbate on foliar uptake was theoretically calculated by Ramge et al. (1993), and experimentally demonstrated by Teklemariam and Sparks. (2006). The enzyme nitrate reductase converts NO_3^- to NO_2^- in 30 the cytosol, and NO_2^- is then transported into the plastids where it is further reduced by the enzyme nitrite reductase to ammonium (NH_4^+), the product required for amino acid synthesis (Ammann et al., 1995; Tischner, 2000; Teklemariam and Sparks, 2006). Alternatively, NO_2 can deposit directly onto the leaf cuticles or a leaf-surface water film (Burkhardt and Eiden, 1994). However, foliar uptake of NO_2 has been demonstrated to be controlled primarily by the stomata, with deposition to the leaf surface representing only a small fraction of the total NO_2 flux (Thoene et al., 1991; Gessler et al.,

2000; Chaparro-Suarez et al., 2011). Strong correlations have been observed between NO_2 concentrations, stomatal conductances, and the NO_2 deposition flux, suggesting foliar uptake is mainly controlled by stomatal aperture and internal leaf resistances (Johansson, 1987; Thoene et al., 1991; Rondon et al., 1993; Meixner et al., 1997; Chaparro-Suarez et al., 2011; Breuninger et al., 2013).

5 Despite the large body of research that exists on the leaf-level deposition of NO_2 to vegetation, discrepancies still exist of NO_2 exchange rates and the role of mesophilic processes. Many laboratory experiments have failed to measure uptake rates necessary to describe the observed 20–50% reduction of soil-emitted NO_x (Hanson and Lindberg, 1991; Breuninger et al., 2013), despite the many modeling studies that have suggested dry deposition makes up most of this reduction (Jacob and Wofsy, 1990; Yienger and Levy, 1995; Ganzeveld et al., 2002a; Geddes and Murphy, 2014). Photolysis 10 gradients and reaction of NO_x to form higher nitrogen oxides could account for a large fraction of this reduction in soil NO_x , as has been suggested by Min et al. (2014, 2012), but the relative importance of dry deposition processes versus in-canopy chemical transformations is still a matter of considerable uncertainty (Lerdau et al., 2000; Ganzeveld et al., 2002a). Another controversy is the existence of a compensation point—a concentration below which leaves would instead act as a source of NO_x . Compensation points of 0.1–3.2 ppb NO_2 have been observed in a number of laboratory chamber studies, suggesting 15 trees instead may serve as a large source of NO_x in forests (Johansson, 1987; Rondon et al., 1993; Hereid and Monson, 2001; Sparks et al., 2001; Teklemariam and Sparks, 2006). Emission of NO at these low NO_x mixing ratios has also been detected in laboratory chamber studies (Wildt et al., 1997; Hereid and Monson, 2001). More recent laboratory studies of leaf level deposition have, however, questioned the existence of a compensation point (Chaparro-Suarez et al., 2011; Breuninger et al., 2013). Most observations of NO_x canopy fluxes and atmospheric models predict or assume substantial NO_x deposition at 20 concentrations as low as 0.1 ppb, typical of NO_x mixing ratios in remote areas (Jacob and Wofsy, 1990; Wang and Leuning, 1998; Lerdau et al., 2000; Sparks et al., 2001; Wolfe and Thornton, 2011; Min et al., 2012; Geddes and Murphy, 2014). However, some modeling studies have suggested that an NO_2 compensation point is necessary to describe (Seok et al., 2013), or has only a small effect on canopy fluxes in most regions (Ganzeveld et al., 2002a). More research is thus needed on 25 leaf and canopy-level processes to understand the full complexity of the soil-canopy-atmosphere system.

25 To understand the leaf-level processes affecting ecosystem scale atmosphere-biosphere NO_x exchange, we have conducted laboratory experiments measuring NO and NO_2 fluxes to the native California tree species *Quercus agrifolia* (Fig. 1) using a branch enclosure system and direct laser-induced fluorescence (LIF) detection of NO_2 (Fig. 2). With the LIF technique we are able measure NO_x exchange fluxes with high specificity and sensitivity at trace NO_x mixing ratios relevant to forested environments. We investigated the existence of an NO_2 and NO compensation point and the rate of NO_x foliar uptake under controlled conditions. To our knowledge this is the first leaf-level uptake experiment that has been performed 30 on a North American tree species.

2 Materials and methods

2.1 *Quercus agrifolia* samples

NO_x uptake by *Quercus agrifolia* (Coastal Live Oak) was investigated in the laboratory. Three *Quercus agrifolia* individuals were purchased from a local native California plant nursery (Native Here Now Nursery), where the plants were grown from 5 seeds and cuttings collected in Contra Costa County. The tree specimens were grown in a nutrient rich commercial soil mixture (a mixture of Orchard Potting Soil and EB Stone Cactus Mix) at the Jane Grey Research Greenhouse at the University of California, Berkeley. The trees were 2–3 years old when measurements were taken.

2.2 Laser-induced fluorescence detection

NO₂ was measured using Laser-Induced Fluorescence (LIF). A blue diode laser (Z-Laser ZM18H3,) centered at a 10 wavelength of 405 nm was focused into each detection cell and made 20 passes in White multipass optical configuration (Fig. 2b)(Thornton et al., 2000; Fuchs et al., 2009). Upon absorption of a visible photon, NO₂ undergoes a transition from the ²A₁ ground to the ²B₂ excited electronic state. The excited NO₂ molecule is either quenched by collision or emits a red-shifted photon as it relaxes back to ground state (e.g. Thornton et al., 2000). These emitted photons were detected using a 15 red-sensitive photomultiplier tube (PMT) (Hamamatsu H7421-50). To minimize collisional quenching, each detection cell was maintained at a pressure of around 3 torr. Excitation at 405 nm was chosen because it is near the region of maximum absorption in the NO₂ spectrum, and is not subject to interferences from absorption by water vapor or O₃(Matsumoto and 20 Kajii, 2003).

Calibrations were performed every hour by diluting NO (4.97 ppm \pm 5%, Praxair) and NO₂ standard gases (5.08 ppm \pm 5%, Praxair) to 1–10 ppb in humidified (RH \sim 60%) zero air. The limit of detection (LOD) for the detection cells is described as:

$$LOD = \frac{S/N}{m} \sqrt{\frac{2b}{t}} \quad (1)$$

where m is the slope of the calibration curve constructed from standard dilutions, b is the PMT signal at 0 ppb NO or NO₂, S/N is the desired signal to noise ratio, and t is the time of signal averaging. At a S/N of 2 and signal averaging over 5 min, the LOD for detection cells 1–4 was 15 ppt, 4 ppt, 10 ppt, and 30 ppt, respectively. NO₂ in the incoming and outgoing 25 airstreams was measured simultaneously in the first two detection cells. In the second two detection cells, NO was quantitatively converted to NO₂ in the presence of excess ozone, allowing for detection of total NO_x (Fig. 2a). Ozone was produced using an ozone generator (Jelight 600) and flow rates of ozone delivered were adjusted to achieve unity conversion of NO to NO₂.

2.3 Dynamic chamber system

The NO_x flux measurements were performed with a dynamic branch enclosure system, consisting of a thin transparent double-walled Teflon film (FEP) bag (American Duraflim), which transmits 90% of photosynthetically active radiation. The chamber was illuminated by an LED diode array of 430–475 nm and 620–670 nm lights (Apollo Horticulture). This light source was selected because it does not emit wavelengths below 420 nm, where NO_2 dissociates, preventing loss of NO_2 to photodissociation and resultant photochemistry. In order to ensure turbulent mixing and minimal aerodynamic and boundary layer resistances, a Teflon-coated fan was installed inside the inner chamber (Meixner et al., 1997; Pape et al., 2009; Breuninger et al., 2013).

Before experiments with *Quercus agrifolia* individuals, the deposition to an empty chamber was measured and background subtracted from subsequent branch measurements. The measured loss of NO_2 to chamber walls was 5% of the NO_2 mixing ratio flowing into chamber. This corresponded to a maximum loss of 0.4 ppb at 8 ppb NO_2 and minimum loss of 0.05 ppb at 1 ppb NO_2 . Emission of less than 0.05 ppb NO_2 from the Teflon walls was also observed when chamber lights were turned on with 0 ppb NO_2 flowing through the system. It is likely that the chamber walls buffer uptake of NO_2 , but this is a minor effect, as the wall emission observed was a tiny fraction of the measured fluxes.

During measurements, the enclosed branch was exposed to known amounts of either NO_2 or NO mixed with zero air. The inner chamber had an inner diameter of 20 cm, a length of 40 cm, and a total volume of 13 L (American Duraflim 200A Teflon FEP). Flow rates into the inner chamber (Q) during experiments were typically 5 L min^{-1} , creating a residence time in the chamber of 3 min. The outer chamber had an inner diameter of 30 cm and a length of 55 cm (American Duraflim 500C20 Teflon FEP). Zero air at a flow rate of 3 L min^{-1} constantly fumigated the outer bag, serving as a buffer region to ensure the laboratory air, with high mixing ratios of NO_x , did not diffuse into the bag enclosing the branch.

The photosynthetic photon flux density (PPFD) was monitored outside the chamber with a LiCor quantum sensor (LiCor LI-190SA). The flux density measured above the chamber was $1190 \mu\text{mol m}^{-2} \text{ s}^{-1}$, approximately the PPFD for Berkeley, California at noon during the month of October. This is well above the photon flux required to achieve maximal stomatal aperture for broadleaf evergreen trees (von Caemmerer and Farquhar, 1981; Chaparro-Suarez et al., 2011; Breuninger et al., 2013). We confirmed this assumption by covering the lights with a filter to reduce the intensity by 40% and monitoring CO_2 and H_2O exchange. No reduction in the exchange rates of these gases were observed. The relative humidity of air entering the chamber was maintained at 50–65% in all experiments by flowing zero air through a bubbler before mixing with NO_x . Measurements of NO_x exchange fluxes occurred under a light/dark cycle with a photoperiod of 12 hours and a temperature of $26/22 \pm 2^\circ\text{C}$. No change in NO_x uptake was observed when heating the chamber with the lights off or cooling the chamber with the lights on. We therefore expect no significant temperature effects caused by the 4°C difference in temperature between light and dark periods. We also observed a relative humidity increase in the delivered air of about 2% with the lights off, but do not expect this increase to produce any significant changes in NO_x deposition or plant physiology (von Caemmerer and Farquhar, 1981; Chaparro-Suarez et al., 2011).

5 Exchange of CO₂ and H₂O with the leaves were monitored with a LiCor-6262 H₂O/CO₂ analyzer operating in differential mode. Flows of 0.1 L min⁻¹ of air entering and exiting the chamber were diverted to the LiCor analyzer to measure the CO₂ assimilation and transpiration rates. To measure the CO₂ content and relative humidity of air delivered to the chamber, 0.5 L min⁻¹ of the humidified zero air/NO_x mixture was diverted to a second external 1.5 L cuvette. The temperature and relative humidity of air entering the chamber were measured with a temperature and relative humidity module in the external cuvette (TE Connectivity HTM2500LF). The CO₂ mixing ratios in the external chamber were monitored with a Vaisala CarboCap GMP343 sensor.

2.4 NO_x flux densities

The leaf-level exchange flux of NO or NO₂ (F_{NO_x}) was calculated according to Eq. 2:

$$10 \quad F_{NO_x} = \frac{Q \cdot (C_0 - C_i)}{A} \quad (2)$$

where Q is the flow rate (m³ s⁻¹), A is the enclosed leaf area (m²), C_0 is the concentration leaving the chamber, and C_i is the concentration entering the chamber (nmol m⁻³). The calculated flux is related to a deposition velocity (Vd_{NO_x}) by Eq. 3:

$$F_{NO_x} = -Vd_{NO_x} \cdot (C_0 + C_{comp}) \quad (3)$$

where C_{comp} is the compensation point, the concentration of NO₂ below which the tree would instead act as a source of NO_x.

15 The deposition velocities were calculated through weighted least squares regression of calculated fluxes and outlet NO_x concentrations (C_o). The absolute value of the slope of the regression line was equal to the deposition velocity, with the x-intercept representing the compensation point concentration. The precision error in the NO_x exchange flux (σ_F) was calculated through propagation of the error in the inlet (σ_{C_i}) and outlet (σ_{C_o}) concentrations (Eq. 4).

$$16 \quad \sigma_F = \frac{Q}{A} \sqrt{\sigma_{C_i}^2 + \sigma_{C_o}^2} \quad (4)$$

20 σ_{C_i} and σ_{C_o} were estimated as the larger of the error in the calibration slopes and the standard deviation of the 5 min signal average. From observations in daily deviations of the flow rate and error in measured leaf area using the ImageJ software (Schneider et al., 2012), we estimate the error in $\frac{Q}{A}$ to be a maximum of 0.005 cm s⁻¹. This usually was only a minor contribution to the total error in the NO_x exchange flux.

The calculated deposition velocity was used to find the total resistance to deposition, R , via Eq. 5.

$$25 \quad Vd_{NO_x} = \frac{1}{R} \quad (5)$$

The total resistance is described by the canopy stomatal resistance model (Baldocchi et al., 1987) and defined in Eq. 6–7.

$$R = R_a + R_b + R_{leaf} \quad (6)$$

$$R_{leaf} = \left(\frac{1}{R_{cut}} + \frac{1}{R_{st} + R_m} \right)^{-1} \quad (7)$$

30 where R_{leaf} is the total leaf resistance and R_a , R_b , R_{cut} , R_{st} , and R_m are the aerodynamic, boundary layer, cuticular, stomatal, and mesophilic resistances, respectively. The aerodynamic resistance is characterized by the micrometeorology

above a surface and is dependent upon the wind speed and turbulence of air flow. The boundary layer resistance describes the diffusion of a molecule through a shallow boundary of air above a surface and is dependent on microscopic surface properties, diffusivity of the gas species, wind speed, and turbulence of air flow (Balocchi et al., 1987). R_{cut} , R_{st} , and R_m are the resistances associated with deposition to the leaf cuticles or through the stomata, and are dependent upon plant 5 physiology.

The chamber fan, installed to create turbulent mixing, allowed for the assumption that R_a was negligible (Pape et al., 2009; Breuninger et al., 2012). R_b is chamber-specific, and has typically not been measured in previous chamber experiments of NO₂ leaf-level deposition (Chaparro-Suarez et al., 2011; Breuninger et al., 2012; Breuninger et al., 2013). R_b was experimentally measured in this study by placing a tray of activated carbon into the chamber (assumed to have zero 10 surface resistance to deposition of NO₂), and calculating the deposition flux of NO₂. The leaf components to the total deposition resistance were determined through dark and light experiments. During dark experiments, the stomata were closed (confirmed with measurements of CO₂ and H₂O exchange), and the deposition observed was assumed to be entirely driven by deposition to the cuticles.

3 Results

15 3.1 Determination of the boundary resistance R_b

To estimate the chamber boundary layer resistance and test the assumption that $R_b \ll R_{leaf}$, a dish of activated carbon, which theoretically has zero chemical resistance to deposition of NO₂, was placed inside the chamber. The boundary layer resistance was considered to be the only component of the total resistance to deposition. The deposition velocity of NO₂ to activated carbon was measured as $0.52 \pm 0.06 \text{ cm s}^{-1}$, corresponding to a boundary layer resistance to NO₂ deposition of $1.94 \pm 0.02 \text{ s cm}^{-1}$ (Fig. 3). This boundary resistance is approximately double what was measured by Pape et al. (2009)—a 20 reasonable difference given differences in chamber design (Fig. 2). The R_b for NO₂ was scaled with the ratio of diffusivities of NO₂ and NO in air to obtain the resistance to deposition of NO of $2.59 \pm 0.03 \text{ s cm}^{-1}$. However, with a branch enclosed inside the chamber, the effective boundary resistance to deposition will likely be reduced, as the surface roughness and surface area for deposition is increased (Galbally and Roy, 1980; Pape et al., 2009). The boundary resistances presented 25 above thus serve as an upper limit for R_b with vegetation inside the chamber.

The boundary resistance was also estimated in an additional experiment (not shown) in which a de-ionized water-soaked Whatman No. 1 filter paper was placed inside the chamber and the evaporation of water vapor into the chamber filled with dry zero air was measured. The emission flux of water vapor from the filter paper was calculated in a similar manner to that of NO_x deposition flux (Eq. 2). The conductance to water vapor was then calculated via:

$$30 \quad \frac{Q \cdot (P_{H_2O})}{A} = g_w (P_{sat} - P_{H_2O}) \quad (8)$$

where P_{H_2O} is the partial pressure of water vapor inside the chamber, P_{sat} is the saturation vapor pressure at the temperature in the chamber, and g_w is the conductance to water vapor. The measured conductance to water vapor was scaled with the ratio of diffusivities of NO_2 to water vapor (D_{NO_2}/D_{H_2O}) and inverted to find the NO_2 boundary layer resistance:

$$R_b = \frac{D_{H_2O}}{D_{\text{NO}_2}} \frac{1}{g_w} \quad (9)$$

5 The boundary resistance to NO_2 deposition by this method was found to be 2 s cm^{-1} , essentially identical to the measurement on the activated-carbon.

3.2 NO_x deposition velocity and compensation point concentration

The deposition velocities and compensation points were respectively calculated as the slope and x-axis intercept of the regression line between NO_x exchange flux and chamber NO_x concentrations (Fig. 4). The detection limit was the dominant

10 source of error in the estimation of the NO exchange flux and compensation point. The large relative uncertainties in NO flux measurements were caused by the much slower deposition of NO compared with that of NO_2 , inhibiting our ability to observe the very small changes between the NO concentration in the chamber and the incoming airstream (Fig. 4). Additional uncertainty in NO_2 flux measurements because of enhanced quenching of NO_2 by water vapor should be minimal, as calibrations and measurements were performed at equivalent relative humidities. However, transpiration of the enclosed 15 leaves caused the absolute humidity within chamber to be enhanced by 0.3–0.5% relative to the incoming airstream. We expect this to result in a maximum error in calculated NO_2 mixing ratios of 1–1.75% (Thornton et al., 2000), resulting in maximum errors in the calculated fluxes and deposition velocities of 2% and 4%, respectively. This 4% error in the calculated deposition velocity during lights-on experiments is somewhat less than the uncertainty of the linear fit (Fig. 4).

Correlation coefficients, deposition velocities, compensation points, and statistical testing of the compensation point 20 for NO_2 and NO deposition are shown in Table 1 and Table 2, respectively, and were calculated according to Breuninger et al. (2013). For NO_2 experiments, only one dark and one light experiment with *Quercus agrifolia* 1, was found to have a statistically significant ($\alpha = 0.05$) non-zero intersection with the x-axis (Table 1). The range of C_{comp} measured was -0.02–0.300 ppb NO_2 , with probabilities of $C_{comp} = 0$ ranging from 10.3–91.6% (excluding the two *Quercus agrifolia* 1 experiments) (Table 1). Conversely, all three *Quercus agrifolia* individuals during all dark and light NO deposition 25 experiments demonstrated compensation points significantly above zero, ranging from 0.74–3.8 ppb NO. The average compensation point was calculated as 0.84 ± 0.32 ppb NO during light experiments and 2.4 ± 1.1 ppb NO during dark experiments (Table 2).

Student's t tests, (not shown), demonstrated that deposition velocities and compensation points measured during NO and NO_2 lights on and off experiments were not significantly different (to the $\alpha=0.05$ confidence level) between 30 different *Quercus agrifolia* individuals. Deposition velocities for NO_2 light experiments were between 0.08 and 0.18 cm s^{-1} , with a deposition of $0.123 \pm 0.009 \text{ cm s}^{-1}$ calculated from the regression of all light experiments. Dark experiments resulted in deposition velocities between 0.013 and 0.022 cm s^{-1} , with a deposition velocity of $0.015 \pm 0.001 \text{ cm s}^{-1}$ calculated from

the regression of all dark experiments (Table 1). NO demonstrated much slower deposition, with deposition velocities from all light and dark experiments calculated as $0.012 \pm 0.002 \text{ cm s}^{-1}$ and $0.005 \pm 0.002 \text{ cm s}^{-1}$, respectively (Table 2). Despite the large compensation point measured for NO, the leaf emission fluxes of NO were a maximum of only $0.8 \text{ pmol m}^{-2} \text{ s}^{-1}$ at 0.1 ppb NO, approximately half of the deposition flux measured for NO_2 at 0.1 ppb (Fig. 4). At typical NO_2/NO ratios and 5 gradients measured in forest canopies, the leaf-level NO_2 and NO exchange fluxes measured make dry stomatal deposition to *Quercus agrifolia* a net sink of NO_x within the canopy.

3.3 Resistances to Leaf-level NO_x deposition.

The deposition velocity measured from linear regression of NO_x exchange fluxes and NO_x chamber concentrations is the inverse of the total resistance to deposition (Eq. 6), with R_a assumed to be zero. The total resistance in the chamber is thus:

$$10 \quad R = R_b + \left(\frac{1}{R_{cut}} + \frac{1}{R_s^*} \right)^{-1} \quad (10)$$

where R_s^* is the sum of R_m and R_{st} . The leaf resistance to deposition can then be found by subtracting the boundary layer resistance from the total resistance. Total leaf resistances, R_{leaf} , were calculated using the boundary layer resistances for NO_2 and NO of $1.94 \pm 0.02 \text{ s cm}^{-1}$ and $2.59 \pm 0.03 \text{ s cm}^{-1}$, respectively. During the dark experiments, R_{leaf} is equal to R_{cut} , and the deposition velocity measured was estimated as the inverse of the sum of the boundary and cuticular resistance. After 15 calculation of R_{cut} from dark experiments, the sum of the stomatal and mesophilic contributions (R_s^*) to the total leaf resistance was determined. The boundary, cuticular, and R_s^* are shown in Table 3. It should be noted that since the reported R_b is the maximum possible boundary resistance, the reported R_{cut} and R_s^* are lower limits. If we were to assume the chamber boundary resistance with the branch enclosed is insignificant ($\sim 0 \text{ s cm}^{-1}$), this would introduce maximum systematic 30% and 3% errors to the calculated $\text{NO}_2 R_s^*$ and R_{cut} , respectively (giving an R_s^* of $9.2 \pm 0.9 \text{ s cm}^{-1}$ and an R_{cut} of $67 \pm 8 \text{ s cm}^{-1}$). The errors in the calculated NO resistances would be negligible.

It is possible that the stomata were not entirely closed during dark experiments. Evidence exists that nocturnal stomatal conductance can be large enough to allow for transpiration (Dawson et al., 2007), and small (within the range of uncertainty observed for the LICOR-6262) emission of water vapor during dark experiments was measured. However, even if all the deposition during dark experiments was stomatal, this would cause only a 0.5 s cm^{-1} reduction in the calculated R_s^* 25 for NO_2 , less than the uncertainty from the error in the measured deposition velocity ($\sim 10\%$ error). The cuticular resistances reported here during dark experiment are nonetheless atmospherically relevant to nighttime NO_x deposition.

4 Discussion

4.1 NO_x deposition velocities and compensation points

The strong linear dependence between NO₂ fluxes and NO₂ chamber concentrations that we observe is consistent with previous observations that NO₂ exchange is largely driven by NO₂ concentration differences between the atmosphere and gaseous phase of the leaf (Rondon and Granat, 1994; Gessler et al., 2000; Hereid and Monson, 2001; Sparks et al., 2001; Teklemariam and Sparks, 2006; Pape et al., 2009; Chaparro-Suarez et al., 2011; Breuninger et al., 2012). Our measurements of NO₂ stomatal resistance parameters for *Quercus agrifolia* represents a stomatal deposition velocity ($1/R_s^*$) of 0.14 ± 0.02 cm s⁻¹. This value is similar to the range of 0.1–0.15 cm s⁻¹ that Chaparro-Suarez et al. (2011) found for two European oak tree species, *Quercus robur* and *Quercus ilex*. The deposition velocity measured here for *Quercus agrifolia* is also much larger than 0.007–0.042 cm s⁻¹ range found for Norway spruce (*Picea abies*) by Breuninger et al. (2012), but surprisingly comparable, given the differences in plant species, to the 0.12 cm s⁻¹ deposition velocity found for maize (*Zea mays*) by Hereid and Monson (2001). We also find here a NO₂ flux at 5 ppb of 0.2 nmol m⁻² s⁻¹, similar in magnitude to the 0.1 nmol m⁻² s⁻¹, 0.15–1.5 nmol m⁻² s⁻¹, and 0.18 nmol m⁻² s⁻¹ fluxes measured for *Fagus sylvatica* (Gessler et al., 2000), tropical Panamanian native trees (Sparks et al., 2001), and periwinkle (*Catharanthus roseus*) (Teklemariam and Sparks, 2006), respectively.

Resistance parameters reported above for NO deposition to *Quercus agrifolia* represent a stomatal deposition velocity of 0.007 ± 0.002 cm s⁻¹ and cuticular deposition velocity of 0.005 ± 0.001 cm s⁻¹. This observation of very minor NO uptake—at least an order of magnitude less than that of NO₂ uptake—is also consistent with previous observations (Hanson and Lindberg, 1991; Hereid and Monson, 2001; Teklemariam and Sparks, 2006). We also detected a statistically significant NO compensation point, with low emissions up to 8 pmol m⁻² s⁻¹ observed below 1 ppb. These observations are similar to the 8–14 pmol m⁻² s⁻¹ emission fluxes of NO reported by Hereid and Monson (2001) and Teklemariam and Sparks (2006) at low NO_x concentrations.

No significant NO₂ compensation point was found for our measurements of *Quercus agrifolia* NO_x uptake. Many previous studies have reported NO₂ compensation points, ranging from 0.1–3.0 ppb, implicating trees as a constant source of NO_x in forest ecosystems (Gessler et al., 2000; Hereid and Monson, 2001; Sparks et al., 2001; Teklemariam and Sparks, 2006). Our findings of a lack of NO₂ compensation point support field observations and modeling studies that have recognized NO₂ dry deposition to vegetation as an important NO_x loss process in forests (Jacob and Wofsy, 1990; Ganzeveld et al., 2002b; Geddes and Murphy, 2014). Our results also support the works of Chaparro-Suarez et al. (2011) and Breuninger et al. (2013), who did not find evidence of an NO₂ compensation point.

The primary difference in our experimental setup, compared to previous dynamic chamber studies that have found a NO₂ compensation point, is the use of a direct NO₂ measurement technique. Measurements of a significant NO₂ compensation point have mostly been obtained using techniques requiring conversion of NO₂, followed by chemiluminescence detection of NO (Gessler et al., 2000; Hereid and Monson, 2001; Sparks et al., 2001; Teklemariam and

Sparks, 2006). Such methods have utilized either non-specific photolytic (Gessler et al., 2000; Hereid and Monson, 2001), luminol (Sparks et al., 2001), or catalytic conversion (Teklemariam and Sparks, 2006) techniques, which may have also resulted in the conversion of PAN, HONO, HNO₃, and other organic nitrates, as well as interferences from alkene + ozone reactions (Carter et al., 2005; Reed et al., 2016). If any of these interfering compounds are not excluded from the chamber system, outgas from the chamber, or form from reactions of biogenic emissions, this would cause an enhancement in observed NO₂ compensation point, and a suppression of observed deposition velocity. Our measurements of NO₂ mixing ratios also demonstrate a much higher degree of precision, due largely to a lower detection limit, than comparable experiments with specific photolytic conversion and chemiluminescence measurement of NO₂ (Chaparro-Suarez et al., 2011; Breuninger et al., 2012; Breuninger et al., 2013). Additionally, previous chamber measurements have sometimes employed chamber setups that would let in a substantial amount of UV light, yet did not exclude photochemical reactions between NO₂, NO, and O₃. Such corrections are excluded here because of our use of chamber lights with only wavelengths above 420 nm. To avoid this issue, other experiments have instead involved a setup including a simultaneously measured blank chamber, which would theoretically allow for correction for any reactions resulting from photolysis of NO₂, O₂, or O₃ (Gessler et al., 2000; Hereid and Monson, 2001). Such corrections might be complicated by secondary chemistry not present in our experiments.

4.2 Implication for canopy NO_x loss

Resistance parameters reported above (Table 3) were used in a 1-D seven-layer multibox model representing chemical reactions, vertical transport, and leaf-level processes scaled to the canopy level to assess the impacts of NO_x deposition velocities on the NO_x lifetime and fluxes. The model is constructed in a manner similar to Wolfe and Thornton (2011). Details will be presented elsewhere. The 1-D model was run for meteorological conditions representing the native habitat of *Quercus agrifolia* and two different leaf area indices (LAI), approximately representing the lower and upper limits of LAI found in California oak woodlands. As shown in Fig. 5a and 5b, the model predicts NO_x deposition to *Q. agrifolia* accounts for 3%–7% of the total NO_x loss within the boundary layer if the only source of NO_x is emissions from the soil. This represents a total NO_x lifetime of 7–7.5 hours in the boundary layer, and a lifetime to deposition of 4–11 days in the boundary layer and 0.5–1.2 hours below the canopy. Under these scenarios approximately 15–30% of soil-emitted NO_x is removed in the canopy (Fig. 6)—on the lower end of the range of 25–80% reduction observed in field studies (Jacob and Wofsy, 1990; Lerdau et al., 2000; Ganzeveld et al., 2002a; Min et al., 2014).

The coastal regions of California where *Q. agrifolia* is found frequently experience much higher NO_x mixing ratios of 10–50 ppb. This is particularly important for oak woodlands of the San Francisco Bay and near Los Angeles areas, where anthropogenic emissions from nearby urban centers are the majority of the NO_x source. To account for this extra NO_x source, additional model runs were done with an added term accounting for NO_x advection from a more concentrated upwind source ($C_{NO_x(adv)}$), with advection treated as a simple mixing process:

$$\left(\frac{dc_{NO_x}}{dt} \right) = -k_{mix} (C_{NO_x} - C_{NO_x(adv)}) \quad (11)$$

where $k_{mix} = 0.3 \text{ h}^{-1}$ and $C_{NO_x(adv)}$ is 10 ppb.

In this case, deposition to *Q. agrifolia* could account for 10–22% of the total NO_x loss in the boundary layer (Fig. 5c,d), representing a lifetime to deposition of 5–14 days in the boundary layer and a total NO_x lifetime of 28–33 hours. Deposition 5 in this higher NO_x scenario decreased the total NO_x lifetime by 3–8 hours, compared with a no-deposition case.

5 Conclusions

This work constitutes the first measurements of NO_2 and NO foliar deposition resistance parameters for a North American tree species. We report observations of leaf-level resistances to NO_2 and NO deposition, corresponding to total deposition velocities of NO_2 and NO of $0.123 \pm 0.007 \text{ cm s}^{-1}$ and $0.012 \pm 0.002 \text{ cm s}^{-1}$ in the light and $0.015 \pm 0.001 \text{ cm s}^{-1}$ 10 and $0.005 \pm 0.002 \text{ cm s}^{-1}$ in the dark, respectively. No compensation point was observed for NO_2 , but compensation points of 0.74–3.8 ppb were recorded for NO. The magnitude of NO emission below the compensation point was significantly less than the magnitude of NO_2 uptake in the same concentration range, making *Q. agrifolia* an overall large net sink of NO_x . The observed deposition is large enough to explain canopy reduction factors observed in canopy-level studies, but is at the lower 15 end of estimated global CRFs. The results of the 1-D multibox model demonstrate that the deposition observed accounts for 5–20% of NO_x removal with a NO_x lifetime to deposition of 0.5–1.2 hours beneath the canopy of a California oak woodland. We show that foliar deposition of NO_x represents a significant removal mechanism of NO_x and can have a large impact on NO_x mixing ratios and fluxes in such ecosystems. Further investigations of NO_2 deposition to a larger variety of plant species under a range of environmental conditions are needed to accurately understand the global impacts of NO_2 deposition across diverse ecosystems.

20 *Acknowledgements.* We would like to thank the two reviewers for their constructive comments. We also wish to gratefully acknowledge financial support from the National Science Foundation (NSF, AGS-1352972). Additional support was provided by a NSF Graduate Research Fellowship to Erin R. Delaria.

References

Ammann, M., vonBallmoos, P., Stalder, M., Suter, M., and Brunold, C.: Uptake and assimilation of atmospheric NO₂-N by spruce needles (*Picea abies*): A field study, *Water Air Soil Poll*, 85, 1497-1502, <https://doi.org/10.1007/Bf00477193>, 1995.

5

Baldocchi, D. D., Hicks, B. B., and Camara, P.: A Canopy Stomatal-Resistance Model for Gaseous Deposition to Vegetated Surfaces, *Atmos Environ*, 21, 91-101, [https://doi.org/10.1016/0004-6981\(87\)90274-5](https://doi.org/10.1016/0004-6981(87)90274-5), 1987.

Breuninger, C., Oswald, R., Kesselmeier, J., and Meixner, F. X.: The dynamic chamber method: trace gas exchange fluxes (NO, NO₂, O₃) between plants and the atmosphere in the laboratory and in the field, *Atmos Meas Tech*, 5, 955-989, <https://doi.org/10.5194/amt-5-955-2012>, 2012.

10

Breuninger, C., Meixner, F. X., and Kesselmeier, J.: Field investigations of nitrogen dioxide (NO₂) exchange between plants and the atmosphere, *Atmos Chem Phys*, 13, 773-790, <https://doi.org/10.5194/acp-13-773-2013>, 2013.

Burkhardt, J., and Eiden, R.: Thin Water Films on Coniferous Needles, *Atmos Environ*, 28, 2001-2011, [https://doi.org/10.1016/1352-2310\(94\)90469-3](https://doi.org/10.1016/1352-2310(94)90469-3), 1994.

15

Caranto, J. D., and Lancaster, K. M.: Nitric oxide is an obligate bacterial nitrification intermediate produced by hydroxylamine oxidoreductase, *P Natl Acad Sci USA*, 114, 8217-8222, <https://doi.org/10.1073/pnas.1704504114>, 2017.

Carter, W. P. L., Cocker, D. R., Fitz, D. R., Malkina, I. L., Bumiller, K., Sauer, C. G., Pisano, J. T., Bufalino, C., and Song, C.: A new environmental chamber for evaluation of gas-phase chemical mechanisms and secondary aerosol formation, *Atmos Environ*, 39, 7768-7788, <https://doi.org/10.1016/j.atmosenv.2005.08.040>, 2005.

20

Chaparro-Suarez, I. G., Meixner, F. X., and Kesselmeier, J.: Nitrogen dioxide (NO₂) uptake by vegetation controlled by atmospheric concentrations and plant stomatal aperture, *Atmos Environ*, 45, 5742-5750, <https://doi.org/10.1016/j.atmosenv.2011.07.021>, 2011.

Crutzen, P. J.: Role of NO and NO₂ in the Chemistry of the Troposphere and Stratosphere, *Annu Rev Earth Pl Sc*, 7, 443-472, <https://doi.org/10.1146/annurev.ea.07.050179.002303>, 1979.

25

Dawson, T. E., Burgess, S. S. O., Tu, K. P., Oliveira, R. S., Santiago, L. S., Fisher, J. B., Simonin, K. A., and Ambrose, A. R.: Nighttime transpiration in woody plants from contrasting ecosystems, *Tree Physiol*, 27, 561-575, <https://doi.org/10.1093/treephys/27.4.561>, 2007.

Fuchs, H., Dube, W. P., Lerner, B. M., Wagner, N. L., Williams, E. J., and Brown, S. S.: A Sensitive and Versatile Detector for Atmospheric NO₂ and NO_x Based on Blue Diode Laser Cavity Ring-Down Spectroscopy, *Environ Sci Technol*, 43, 7831-7836, <https://doi.org/10.1021/es902067h>, 2009.

30

Galbally, I. E., and Roy, C. R.: Destruction of Ozone at the Earths Surface, *Q J Roy Meteor Soc*, 106, 599-620, <https://doi.org/10.1002/qj.49710644915>, 1980.

Ganzeveld, L. N., Lelieveld, J., Dentener, F. J., Krol, M. C., Bouwman, A. J., and Roelofs, G. J.: Global soil-biogenic NO_x emissions and the role of canopy processes, *J Geophys Res-Atmos*, 107, Artn 4298, <https://doi.org/10.1029/2001jd001289>, 2002a.

Ganzeveld, L. N., Lelieveld, J., Dentener, F. J., Krol, M. C., and Roelofs, G. J.: Atmosphere-biosphere trace gas exchanges simulated with a single-column model, *J Geophys Res-Atmos*, 107, Artn 4297, <https://doi.org/10.1029/2001jd000684>, 2002b.

Geddes, J. A., and Murphy, J. G.: Observations of reactive nitrogen oxide fluxes by eddy covariance above two midlatitude North American mixed hardwood forests, *Atmos Chem Phys*, 14, 2939-2957, <https://doi.org/10.5194/acp-14-2939-2014>, 2014.

10 Gessler, A., Rienks, M., and Rennenberg, H.: NH₃ and NO₂ fluxes between beech trees and the atmosphere - correlation with climatic and physiological parameters, *New Phytol*, 147, 539-560, <https://doi.org/10.1046/j.1469-8137.2000.00712.x>, 2000.

Hanson, P. J., and Lindberg, S. E.: Dry Deposition of Reactive Nitrogen-Compounds - a Review of Leaf, Canopy and Non-Foliar Measurements, *Atmos Environ a-Gen*, 25, 1615-1634, [https://doi.org/10.1016/0960-1686\(91\)90020-8](https://doi.org/10.1016/0960-1686(91)90020-8), 1991.

15 Hereid, D. P., and Monson, R. K.: Nitrogen oxide fluxes between corn (*Zea mays L.*) leaves and the atmosphere, *Atmos Environ*, 35, 975-983, [https://doi.org/10.1016/S1352-2310\(00\)00342-3](https://doi.org/10.1016/S1352-2310(00)00342-3), 2001.

Jacob, D. J., and Wofsy, S. C.: Budgets of Reactive Nitrogen, Hydrocarbons, and Ozone over the Amazon-Forest during the Wet Season, *J Geophys Res-Atmos*, 95, 16737-16754, <https://doi.org/10.1029/JD095iD10p16737>, 1990.

Johansson, C.: Pine forest: a negligible sink for atmospheric NO_x in rural Sweden, *Tellus B*, 39B, 426-438, 1987.

20 Lee, Y. N., and Schwartz, S. E.: Evaluation of the Rate of Uptake of Nitrogen-Dioxide by Atmospheric and Surface Liquid Water, *J Geophys Res-Oceans*, 86, 1971-1983, <https://doi.org/10.1029/JC086iC12p11971>, 1981a.

Lee, Y. N., and Schwartz, S. E.: Reaction-Kinetics of Nitrogen-Dioxide with Liquid Water at Low Partial-Pressure, *J Phys Chem-Us*, 85, 840-848, <https://doi.org/10.1021/j150607a022>, 1981b.

25 Lerdau, M. T., Munger, L. J., and Jacob, D. J.: Atmospheric chemistry - The NO₂ flux conundrum, *Science*, 289, 2291-+, <https://doi.org/10.1126/science.289.5488.2291>, 2000.

Matsumoto, J., and Kajii, Y.: Improved analyzer for nitrogen dioxide by laser-induced fluorescence technique, *Atmos Environ*, 37, 4847-4851, <https://doi.org/10.1016/j.atmosenv.2003.08.023>, 2003.

Mckenney, D. J., Shuttleworth, K. F., Vriesacker, J. R., and Findlay, W. I.: Production and Loss of Nitric-Oxide from Denitrification in Anaerobic Brookston Clay, *Appl Environ Microb*, 43, 534-541, 1982.

30 Meixner, F. X., Fickinger, T., Marufu, L., Serca, D. N., F. J., Makina, E., Mukurumbira, L., and Andreae, M. O.: Preliminary results on nitric oxide emission from a southern African savanna ecosystem, *Nutrient Cycling in Agroecosystems*, 48, 123-138, 1997.

Min, K. E., Pusede, S. E., Browne, E. C., LaFranchi, B. W., Wooldridge, P. J., Wolfe, G. M., Harrold, S. A., Thornton, J. A., and Cohen, R. C.: Observations of atmosphere-biosphere exchange of total and speciated peroxy nitrates: nitrogen

fluxes and biogenic sources of peroxy nitrates, *Atmos Chem Phys*, 12, 9763-9773, <https://doi.org/10.5194/acp-12-9763-2012>, 2012.

Min, K. E., Pusede, S. E., Browne, E. C., LaFranchi, B. W., Wooldridge, P. J., and Cohen, R. C.: Eddy covariance fluxes and vertical concentration gradient measurements of NO and NO₂ over a ponderosa pine ecosystem: observational evidence for within-canopy chemical removal of NO_x, *Atmos Chem Phys*, 14, 5495-5512, <https://doi.org/10.5194/acp-14-5495-2014>, 2014.

Neubert, A., Kley, D., Wildt, J., Segschneider, H. J., and Forstel, H.: Uptake of NO, NO₂ and O₃ by Sunflower (*Helianthus-Annus* L) and Tobacco Plants (*Nicotiana-Tabacum*-L) - Dependence on Stomatal Conductivity, *Atmos Environ a-Gen*, 27, 2137-2145, [https://doi.org/10.1016/0960-1686\(93\)90043-X](https://doi.org/10.1016/0960-1686(93)90043-X), 1993.

10 Nussbaum, S., Vonballmoos, P., Gfeller, H., Schlunegger, U. P., Fuhrer, J., Rhodes, D., and Brunold, C.: Incorporation of Atmospheric (NO₂)-N-15-Nitrogen into Free Amino-Acids by Norway Spruce *Picea-Abies* (L) Karst, *Oecologia*, 94, 408-414, <https://doi.org/10.1007/Bf00317117>, 1993.

Okano, K., and Totsuka, T.: Absorption of Nitrogen-Dioxide by Sunflower Plants Grown at Various Levels of Nitrate, *New Phytol*, 102, 551-562, <https://doi.org/10.1111/j.1469-8137.1986.tb00831.x>, 1986.

15 Pape, L., Ammann, C., Nyfeler-Brunner, A., Spirig, C., Hens, K., and Meixner, F. X.: An automated dynamic chamber system for surface exchange measurement of non-reactive and reactive trace gases of grassland ecosystems, *Biogeosciences* 6, 405-429, 2009.

Ramge, P., Badeck, F. W., Plochl, M., and Kohlmaier, G. H.: Apoplastic Antioxidants as Decisive Elimination Factors within the Uptake Process of Nitrogen-Dioxide into Leaf Tissues, *New Phytol*, 125, 771-785, <https://doi.org/10.1111/j.1469-8137.1993.tb03927.x>, 1993.

20 Reed, C., Evans, M. J., Di Carlo, P., Lee, J. D., and Carpenter, L. J.: Interferences in photolytic NO₂ measurements: explanation for an apparent missing oxidant?, *Atmos Chem Phys*, 16, 4707-4724, <https://doi.org/10.5194/acp-16-4707-2016>, 2016.

Rogers, H. H., Jeffries, H. E., and Witherspoon, A. M.: Measuring Air Pollutant Uptake by Plants - Nitrogen-Dioxide, *J Environ Qual*, 8, 551-557, <https://doi.org/10.2134/jeq1979.00472425000800040022x>, 1979.

Rondon, A., Johansson, C., and Granat, L.: Dry Deposition of Nitrogen-Dioxide and Ozone to Coniferous Forests, *J Geophys Res-Atmos*, 98, 5159-5172, <https://doi.org/10.1029/92jd02335>, 1993.

Rondon, A., and Granat, L.: Studies on the Dry Deposition of NO₂ to Coniferous Species at Low NO₂ Concentrations, *Tellus B*, 46, 339-352, <https://doi.org/10.1034/j.1600-0889.1994.t01-4-00001.x>, 1994.

30 Schneider, C. A., Rasband, W. S., and Eliceiri, K. W.: NIH Image to ImageJ: 25 years of image analysis, *Nat Methods*, 9, 671-675, <https://doi.org/10.1038/nmeth.2089>, 2012.

Seinfeld, J. H., and Pandis, S. N.: *Atmospheric Chemistry and Physics: From Air Pollution to Global Change*, 2nd ed., Wiley, New York, 2006.

Seok, B., Helming, D., Ganzeveld, L., Williams, M. W., and Vogel, C. S.: Dynamics of nitrogen oxides and ozone above and within a mixed hardwood forest in northern Michigan, *Atmos Chem Phys*, 13, 7301-7320, <https://doi.org/10.5194/acp-13-7301-2013>, 2013.

Sparks, J. P., Monson, R. K., Sparks, K. L., and Lerdau, M.: Leaf uptake of nitrogen dioxide (NO₂) in a tropical wet forest: 5 implications for tropospheric chemistry, *Oecologia*, 127, 214-221, <https://doi.org/10.1007/s004420000594>, 2001.

Teklemariam, T. A., and Sparks, J. P.: Leaf fluxes of NO and NO₂ in four herbaceous plant species: The role of ascorbic acid, *Atmos Environ*, 40, 2235-2244, <https://doi.org/10.1016/j.atmosenv.2005.12.010>, 2006.

Thoene, B., Schroder, P., Papen, H., Egger, A., and Rennenberg, H.: Absorption of Atmospheric NO₂ by Spruce (Picea-Abies L Karst) Trees .1. NO₂ Influx and Its Correlation with Nitrate Reduction, *New Phytol*, 117, 575-585, 10 <https://doi.org/10.1111/j.1469-8137.1991.tb00962.x>, 1991.

Thornton, J. A., Wooldridge, P. J., and Cohen, R. C.: Atmospheric NO₂: In situ laser-induced fluorescence detection at parts per trillion mixing ratios, *Anal Chem*, 72, 528-539, <https://doi.org/10.1021/ac9908905>, 2000.

Tischner, R.: Nitrate uptake and reduction in higher and lower plants, *Plant Cell Environ*, 23, 1005-1024, DOI 15 [10.1046/j.1365-3040.2000.00595.x](https://doi.org/10.1046/j.1365-3040.2000.00595.x), 2000.

Vinken, G. C. M., Boersma, K. F., Maasakkers, J. D., Adon, M., and Martin, R. V.: Worldwide biogenic soil NO_x emissions inferred from OMI NO₂ observations, *Atmos Chem Phys*, 14, 10363-10381, <https://doi.org/10.5194/acp-14-10363-2014>, 2014.

von Caemmerer, S., and Farquhar, G. D.: Some Relationships between the Biochemistry of Photosynthesis and the Gas-Exchange of Leaves, *Planta*, 153, 376-387, 1981.

Wang, Y. P., and Leuning, R.: A two-leaf model for canopy conductance, photosynthesis and partitioning of available energy I: Model description and comparison with a multi-layered model, *Agr Forest Meteorol*, 91, 89-111, 20 [https://doi.org/10.1016/S0168-1923\(98\)00061-6](https://doi.org/10.1016/S0168-1923(98)00061-6), 1998.

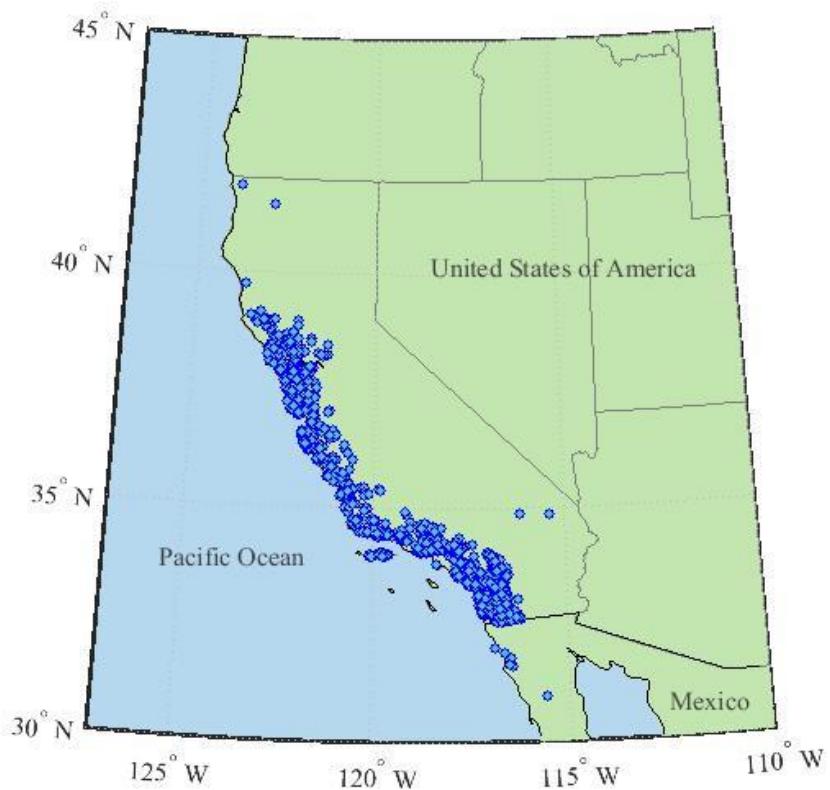
Weber, P., Nussbaum, S., Fuhrer, J., Gfeller, H., Schlunegger, U. P., Brunold, C., and Rennenberg, H.: Uptake of Atmospheric (NO₂)-N-15 and Its Incorporation into Free Amino-Acids in Wheat (Triticum-Aestivum), *Physiol Plantarum*, 94, 71-77, <https://doi.org/10.1034/j.1399-3054.1995.940111.x>, 1995.

Wildt, J., Kley, D., Rockel, A., Rockel, P., and Segschneider, H. J.: Emission of NO from several higher plant species, *J Geophys Res-Atmos*, 102, 5919-5927, <https://doi.org/10.1029/96jd02968>, 1997.

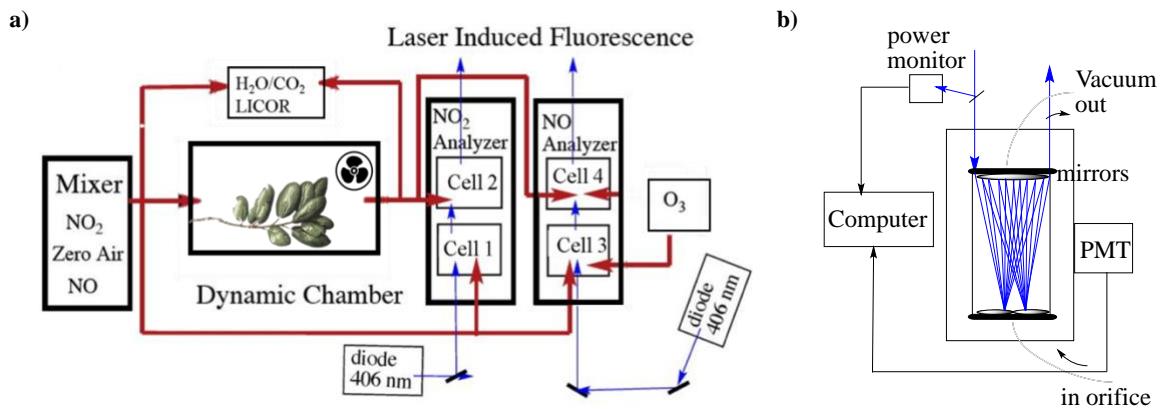
Wolfe, G. M., and Thornton, J. A.: The Chemistry of Atmosphere-Forest Exchange (CAFE) Model - Part 1: Model description and characterization, *Atmos Chem Phys*, 11, 77-101, [10.5194/acp-11-77-2011](https://doi.org/10.5194/acp-11-77-2011), 2011.

Yienger, J. J., and Levy, H.: Empirical-Model of Global Soil-Biogenic NO_x Emissions, *J Geophys Res-Atmos*, 100, 11447-30 11464, <https://doi.org/10.1029/95jd00370>, 1995.

Yoneyama, T., Ito, O., and Engelaar, W. M. H. G.: Uptake, metabolism and distribution of nitrogen in crop plants traced by enriched and natural ¹⁵N: Progress over the last 30 years, *Phytochemistry Reviews*, 2, 121-132, <https://doi.org/10.1023/B:PHYT.0000004198.95836.ad>, 2003.



5 **Figure 1: Species distribution map of *Quercus agrifolia*. Each dot represents an observation of *Q. agrifolia* occurrence. Data provided by the participants of the Consortium of California Herbaria.**



5 **Figure 2: Schematic of the experimental dynamic chamber (a) and laser-induced fluorescence detection (b) setups.**

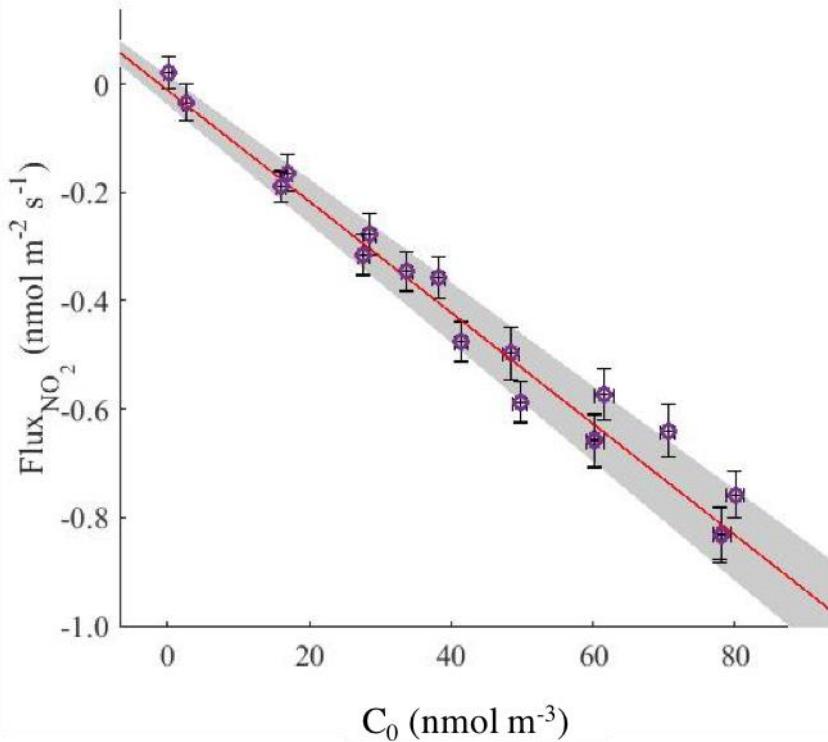


Figure 3: Flux to a 5.1 cm diameter dish filled with activated charcoal. The chemical surface resistance to deposition is approximately zero, so the deposition velocity for deposition of NO_2 to the surface of the charcoal dish is the reciprocal of the boundary layer resistance. The line of best fit is $(0.51 \pm 0.032)C_o$, where C_o is the concentration of NO_2 in the outgoing airstream.

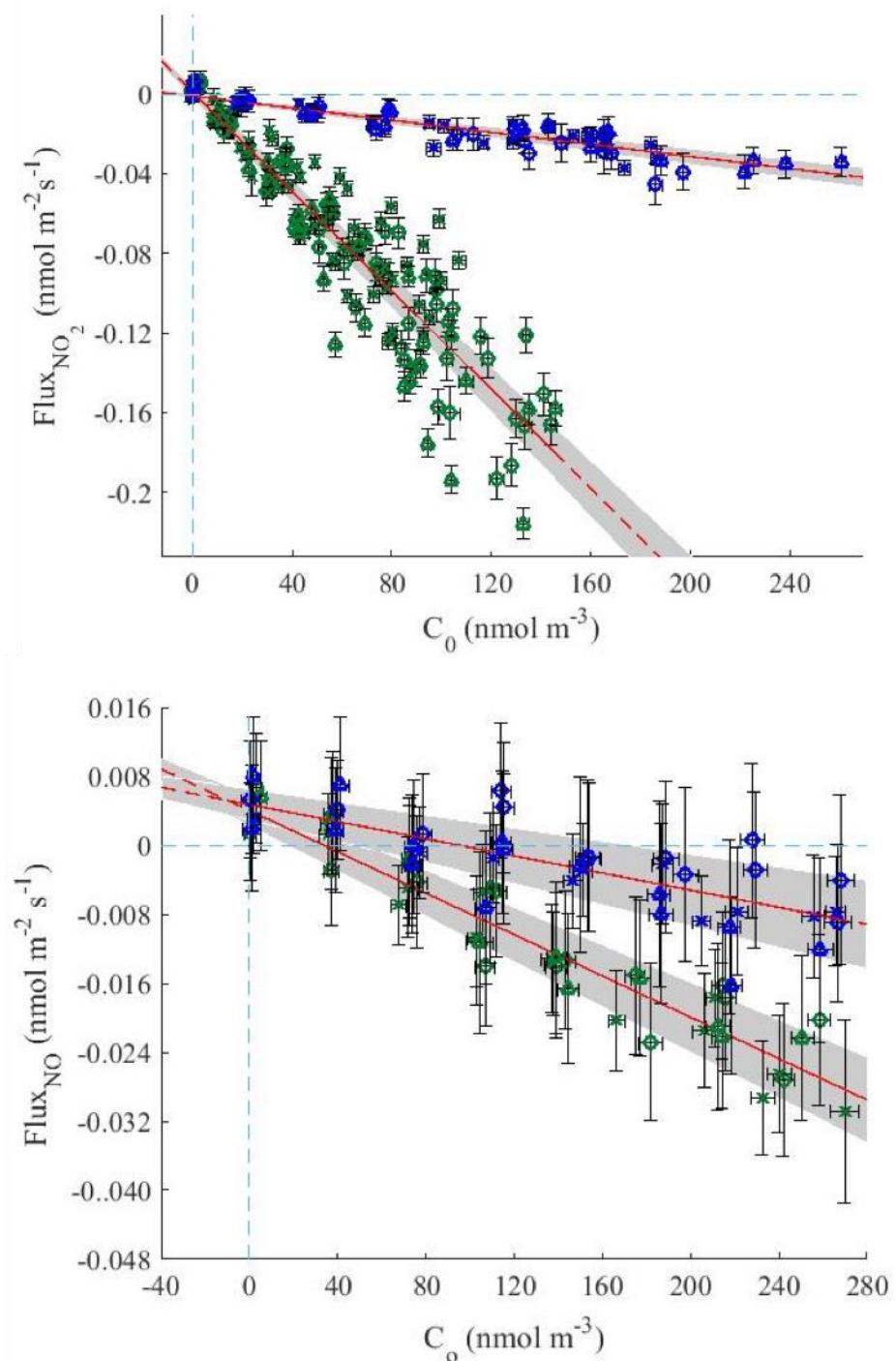


Figure 4: NO_2 (top) and NO (bottom) fluxes versus the outlet concentrations for all *Quercus agrifolia* individuals with the chamber lights on (green) and off (blue). The line of best fit is shown in red and was calculated to minimize the weighted residuals in both the x- and y-axis. The blue dotted line shows where flux is zero. A significantly positive ($\alpha = 0.5$) x-intercept occurs for NO , but not NO_2 experiments.

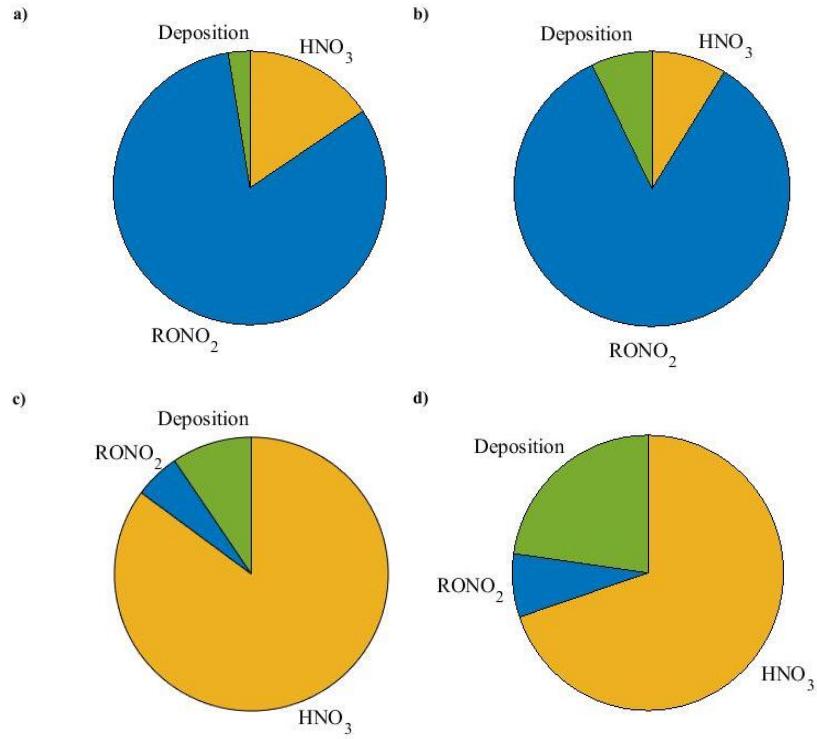
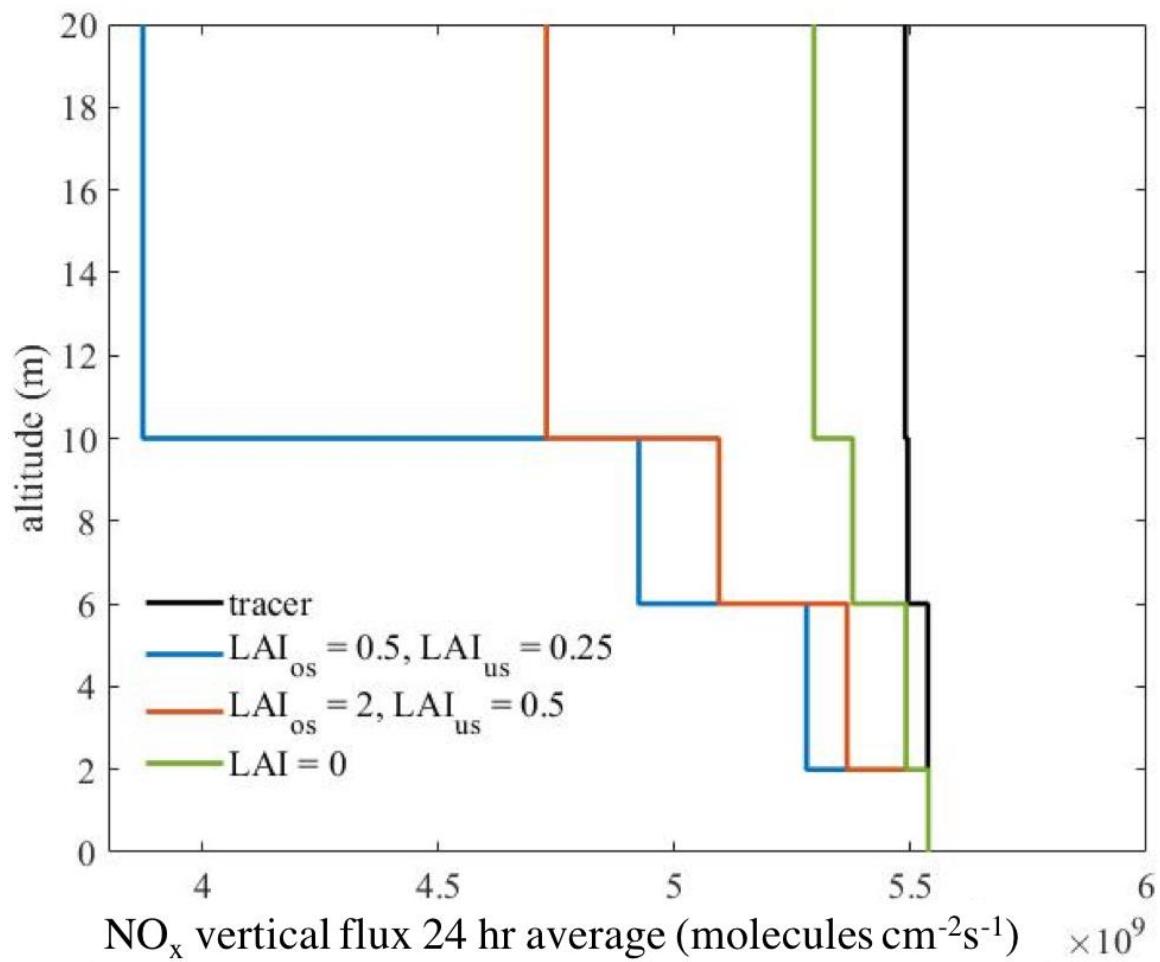


Figure 5: Model predictions of the fraction of NO_x loss to alkyl nitrate formation, nitric acid formation, and deposition in a *Q. agrifolia* woodland. The model was run using scenarios with only soil emissions and LAI of $1 \text{ m}^2/\text{m}^2$ (a), only soil emissions and LAI of $3 \text{ m}^2/\text{m}^2$ (b), $C_{\text{NO}_x(\text{adv})} = 10 \text{ ppb}$ and LAI of $1 \text{ m}^2/\text{m}^2$ (c), and $C_{\text{NO}_x(\text{adv})} = 10 \text{ ppb}$ and LAI of $3 \text{ m}^2/\text{m}^2$ (d).



5

Figure 6: 24 hr average vertical fluxes of NO_x predicted by the 1-D multibox model for a California oak woodland using the leaf resistances measured in this study. Model runs were conducted for a low (red) and high (blue) LAI cases and for a no deposition scenario (green).

Table 1: Parameters of NO₂ bi-variate linear least-square fitting regression analysis

run	N	R ²	[NO ₂] _{comp} (ppb)	P([NO ₂] _{comp} =0) %	V _{dep} cm s ⁻¹
<i>Q.agrifolia</i> 1, light					
1	13	0.979	0.056 ± 0.013	42.7	0.10 ± 0.013
2	13	0.950	0.046 ± 0.19	63.7	0.12 ± 0.023
3	16	0.978	0.099 ± 0.086	3.87	0.15 ± 0.016
4	16	0.958	0.077 ± 0.14	28.7	0.12 ± 0.021
All	58	0.927	0.080 ± 0.10	11.6	0.12 ± 0.012
<i>Q.agrifolia</i> 2, light					
1	16	0.963	0.10 ± 0.12	10.3	0.08 ± 0.011
2	5	0.969	-0.01 ± 0.96	83.8	0.12 ± 0.014
3	9	0.997	0.023 ± 0.032	20.3	0.16 ± 0.011
4	16	0.974	-0.019 ± 0.074	61.9	0.14 ± 0.017
5	15	0.979	0.015 ± 0.082	72.7	0.12 ± 0.014
All	61	0.845	-0.0077 ± 0.091	91.6	0.11 ± 0.014
<i>Q.agrifolia</i> 3, light					
1	11	0.969	0.016 ± 0.18	87.4	0.12 ± 0.024
2	15	0.961	0.074 ± 0.16	39.1	0.18 ± 0.029
3	5	0.990	0.30 ± 0.20	5.9	0.12 ± 0.038
All	31	0.830	0.019 ± 0.064	77.6	0.14 ± 0.029
All <i>Q.agrifolia</i> , light	150	0.885	0.030 ± 0.072	41.3	0.123 ± 0.0092
<i>Q.agrifolia</i> 1, dark					
1	16	0.964	0.056 ± 0.14	0.9*	0.022 ± 0.0034
<i>Q.agrifolia</i> 2, dark					
1	16	0.858	-0.16 ± 0.47	50.8	0.016 ± 0.0050
2	12	0.932	-0.34 ± 0.40	11.8	0.013 ± 0.0038
All	28	0.853	-0.24 ± 0.32	15.6	0.015 ± 0.0030
<i>Q.agrifolia</i> 3, dark					
1	14	0.900	-0.30 ± 0.48	24.1	0.015 ± 0.0042
2	11	0.909	-0.001 ± 0.69	36.7	0.015 ± 0.0057
All	25	0.898	-0.22 ± 0.38	25.3	0.014 ± 0.0029
All <i>Q.agrifolia</i> , dark	69	0.881	-0.16 ± 0.24	12.2	0.015 ± 0.0018

* Significant non-zero compensation point

Table 2: Parameters of NO bi-variate linear least-square fitting regression analysis

run	N	R ²	[NO ₂] _{comp} (ppb)	P([NO ₂] _{comp} =0)	V _{dep}
<i>Q. agrifolia</i> 1					
light	17	0.874	0.74 ± 0.65	3.5*	0.011 ± 0.0032
dark	13	0.699	3.8 ± 2.2	0.52*	0.0040 ± 0.0025
<i>Q. agrifolia</i> 1					
light	14	0.954	0.76 ± 0.49	0.92*	0.013 ± 0.0027
dark	10	0.866	1.7 ± 1.0	1.1*	0.0046 ± 0.0018
<i>Q. agrifolia</i> 1					
light	12	0.936	1.3 ± 0.60	0.17*	0.0123 ± 0.0029
dark	15	0.803	2.0 ± 1.0	2.5*	0.0074 ± 0.0033
All <i>Q. agrifolia</i>					
light	13	0.908	0.84 ± 0.32	<0.01*	0.012 ± 0.0015
dark	13	0.602	2.4 ± 1.1	<0.01*	0.0050 ± 0.0016

*Significant non-zero compensation point

Table 3: Summary of deposition resistance parameters of *Quercus agrifolia*

gas	R_b $s\text{ cm}^{-1}$	R_{cut} $s\text{ cm}^{-1}$	R_s^* $s\text{ cm}^{-1}$
NO ₂	1.94 ± 0.02	65 ± 8	6.9 ± 0.9
NO	2.59 ± 0.03	200 ± 60	140 ± 40