

**Characterization of Aerosol Composition, Aerosol Acidity and Organic Acid Partitioning at an Agriculture-Intensive Rural Southeastern U.S. Site**

Theodora Nah,<sup>1</sup> Hongyu Guo,<sup>1</sup> Amy P. Sullivan,<sup>2</sup> Yunle Chen,<sup>1</sup> David J. Tanner,<sup>1</sup> Athanasios Nenes,<sup>1,3,4,5</sup> Armistead Russell,<sup>6</sup> Nga Lee Ng,<sup>1,3</sup> L. Gregory Huey<sup>1</sup> and Rodney J. Weber<sup>1,\*</sup>

<sup>1</sup>*School of Earth and Atmospheric Sciences, Georgia Institute of Technology, Atlanta, GA, USA*

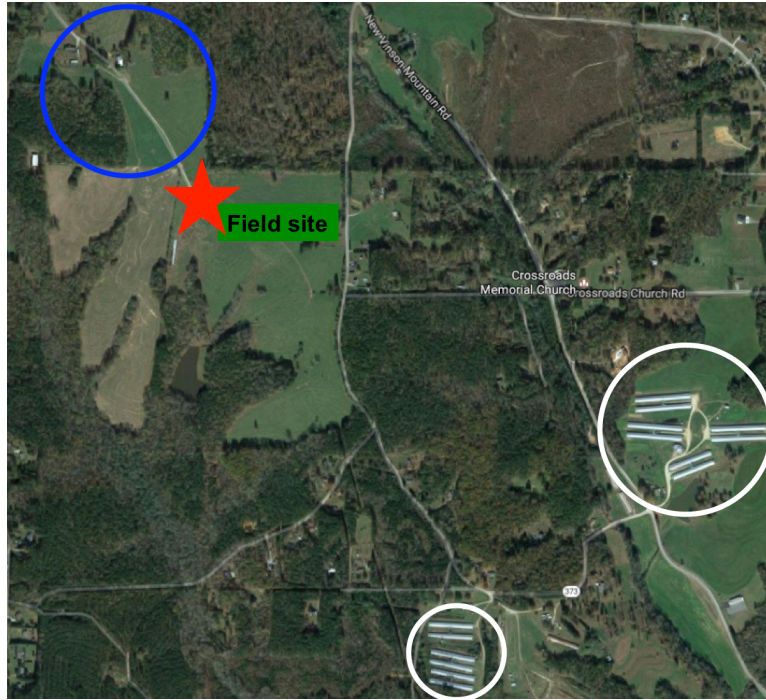
<sup>2</sup>*Department of Atmospheric Science, Colorado State University, Fort Collins, CO, USA*

<sup>3</sup>*School of Chemical and Biomolecular Engineering, Georgia Institute of Technology, Atlanta, GA, USA*

<sup>4</sup>*ICE-HT, Foundation for Research and Technology, Hellas, 26504 Patras, Greece*

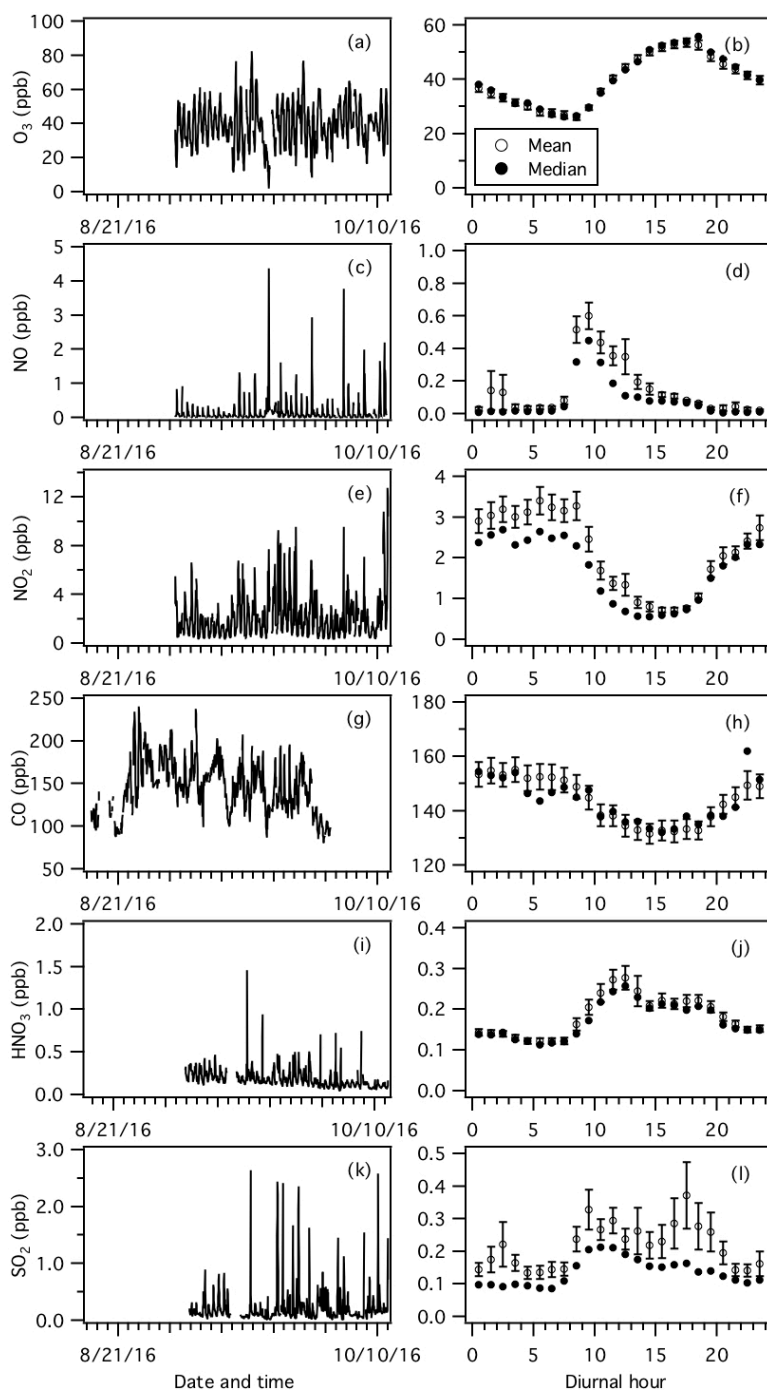
<sup>5</sup>*IERSD, National Observatory of Athens, P. Penteli, 15236, Athens, Greece*

<sup>6</sup>*School of Civil and Environmental Engineering, Georgia Institute of Technology, Atlanta, GA, USA*



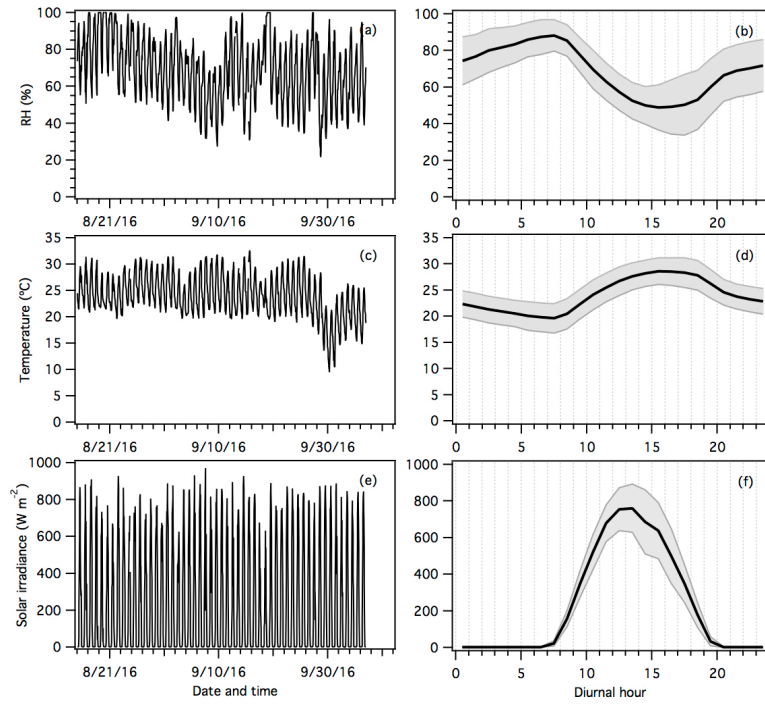
21

22 **Figure S1:** Map of the Yorkville field site (marked by the red diamond) and its surrounding areas  
23 (from Google Maps). Locations of nearby cattle-grazing pastures (north-west direction) and  
24 poultry CAFOs (south-east direction) are marked by the blue and white circles, respectively.

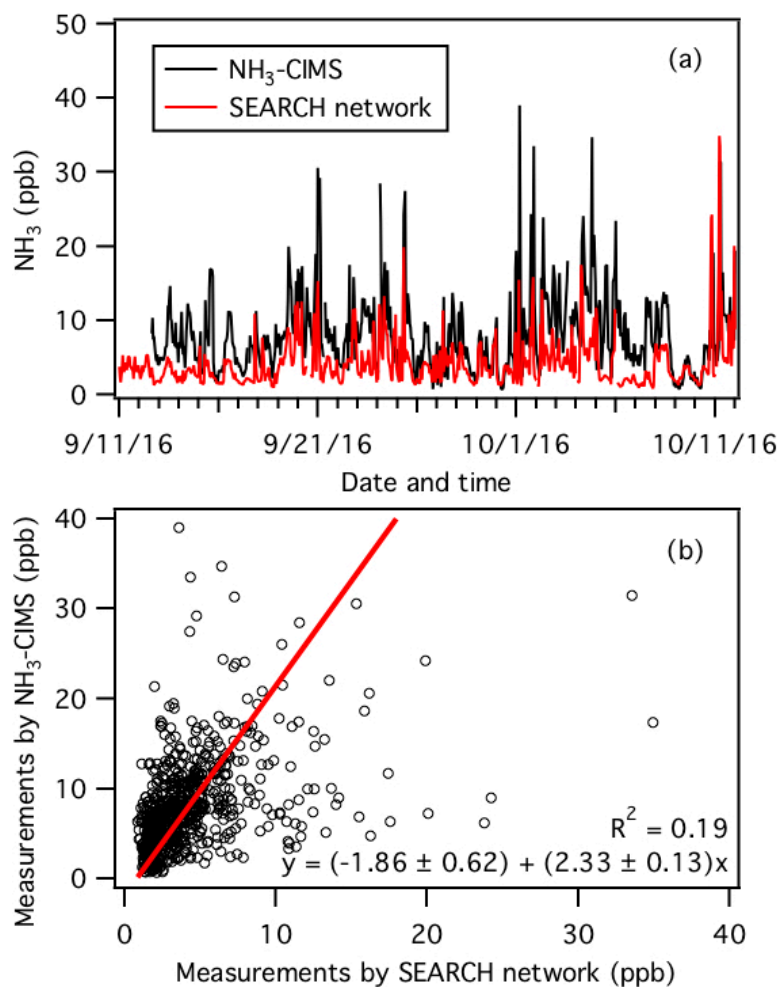


25

26 **Figure S2:** Time series and diurnal profiles of (a and b) O<sub>3</sub>, (c and d) NO, (e and f) NO<sub>2</sub>, (g and  
 27 h) CO, (i and j) HNO<sub>3</sub>, and (k and l) SO<sub>2</sub>. Dates and times displayed are local time. All the  
 28 concentrations represent averages in 1-hour intervals and the standard errors are plotted as error  
 29 bars. O<sub>3</sub>, NO, NO<sub>2</sub> and CO measurements were provided by the SEARCH network. HNO<sub>3</sub> and  
 30 SO<sub>2</sub> were measured by the SF<sub>6</sub>-CIMS.

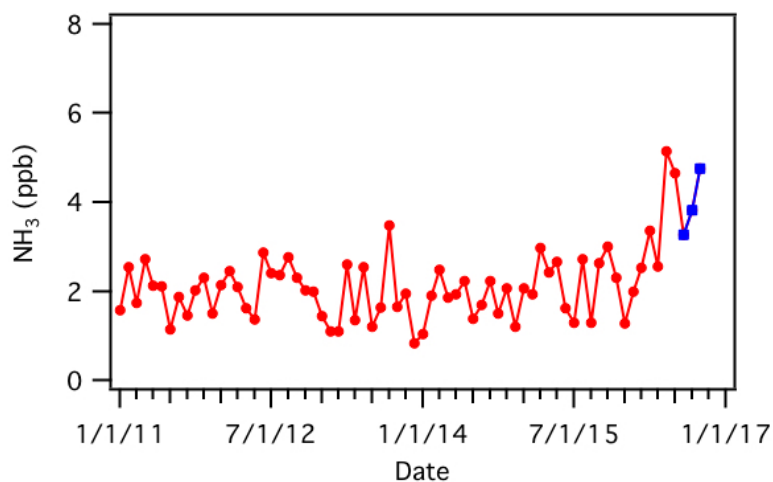


**Figure S3:** Time series and diurnal profiles of (a and b) RH, (c and d) temperature, and (e and f) solar irradiance. Dates and times displayed are local time. In panels b, d and f, the lines within the shaded area represents the average values. The upper and lower boundaries of the shaded areas mark one standard deviation. RH, temperature and solar irradiance measurements were provided by the SEARCH network.

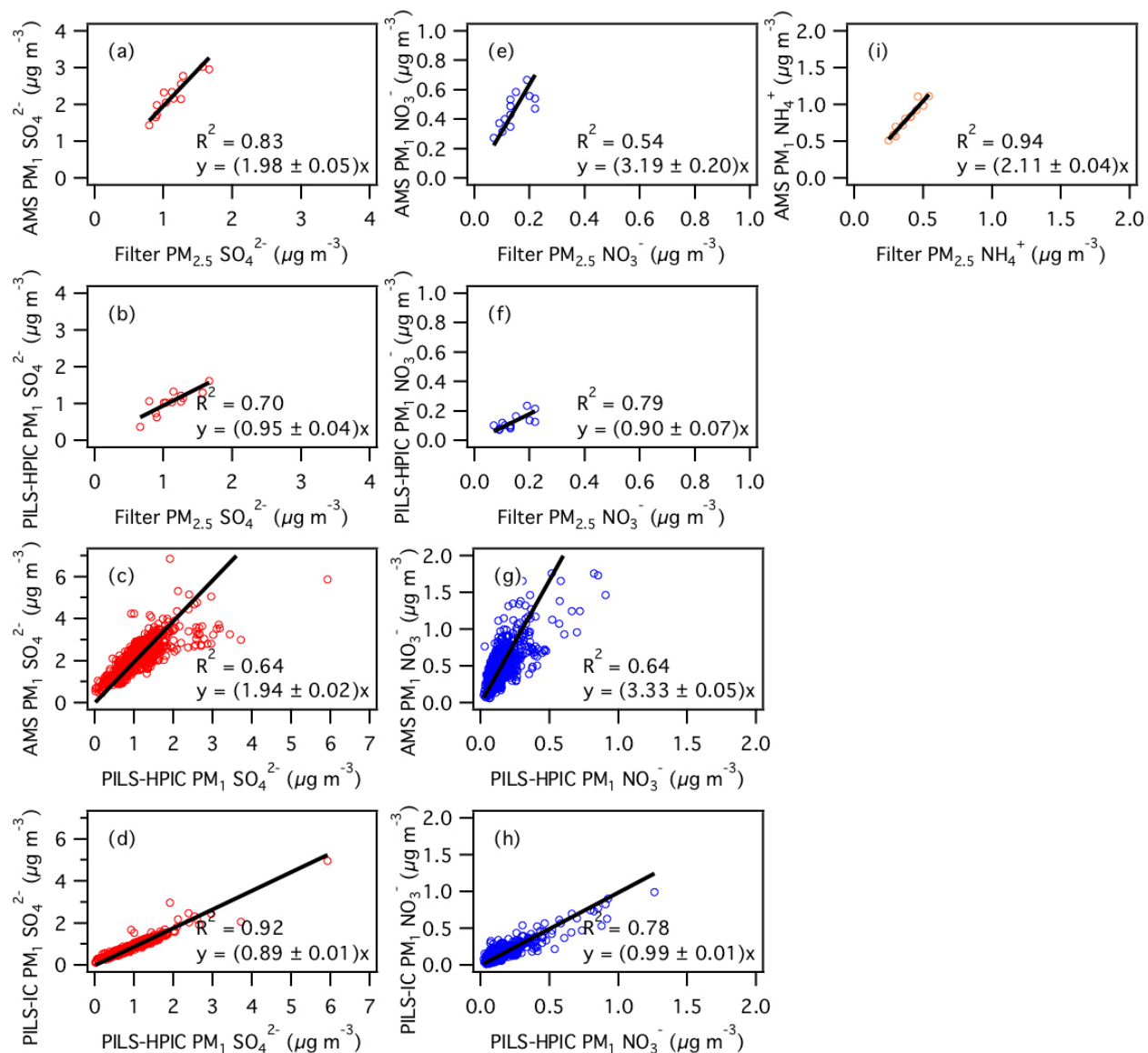


37

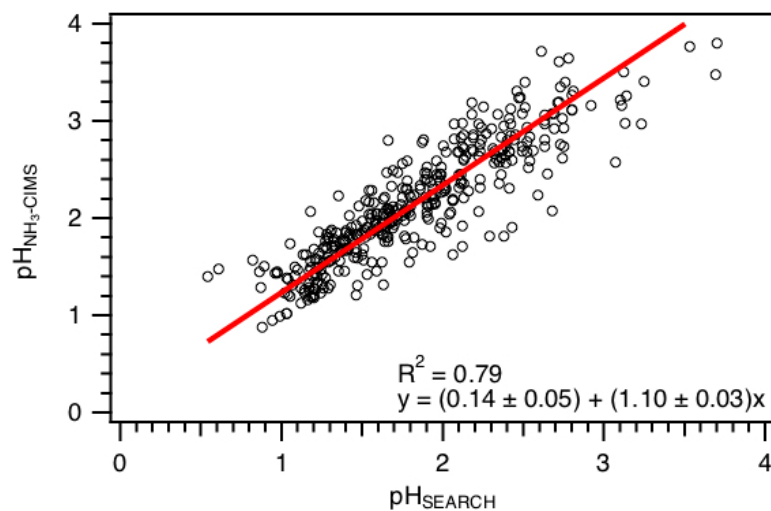
38 **Figure S4:** (a) Time series of  $\text{NH}_3$  concentrations measured by the  $\text{NH}_3$ -CIMS and denuder-based  
 39 instrument operated by the SEARCH network. (b) Comparison of  $\text{NH}_3$  concentrations measured  
 40 by the  $\text{NH}_3$ -CIMS and denuder-based instrument. The red line is the orthogonal distance regression  
 41 fit to the data. All the data are displayed as 1-hour averages.



**Figure S5:** Monthly-averaged NH<sub>3</sub> concentrations at the Yorkville SEARCH monitoring site for 2011 to 2016. These measurements were made using the denuder-based instrument operated by the SEARCH network. Concentrations measured during this study (mid-August to mid-October 2016) are shown in blue.



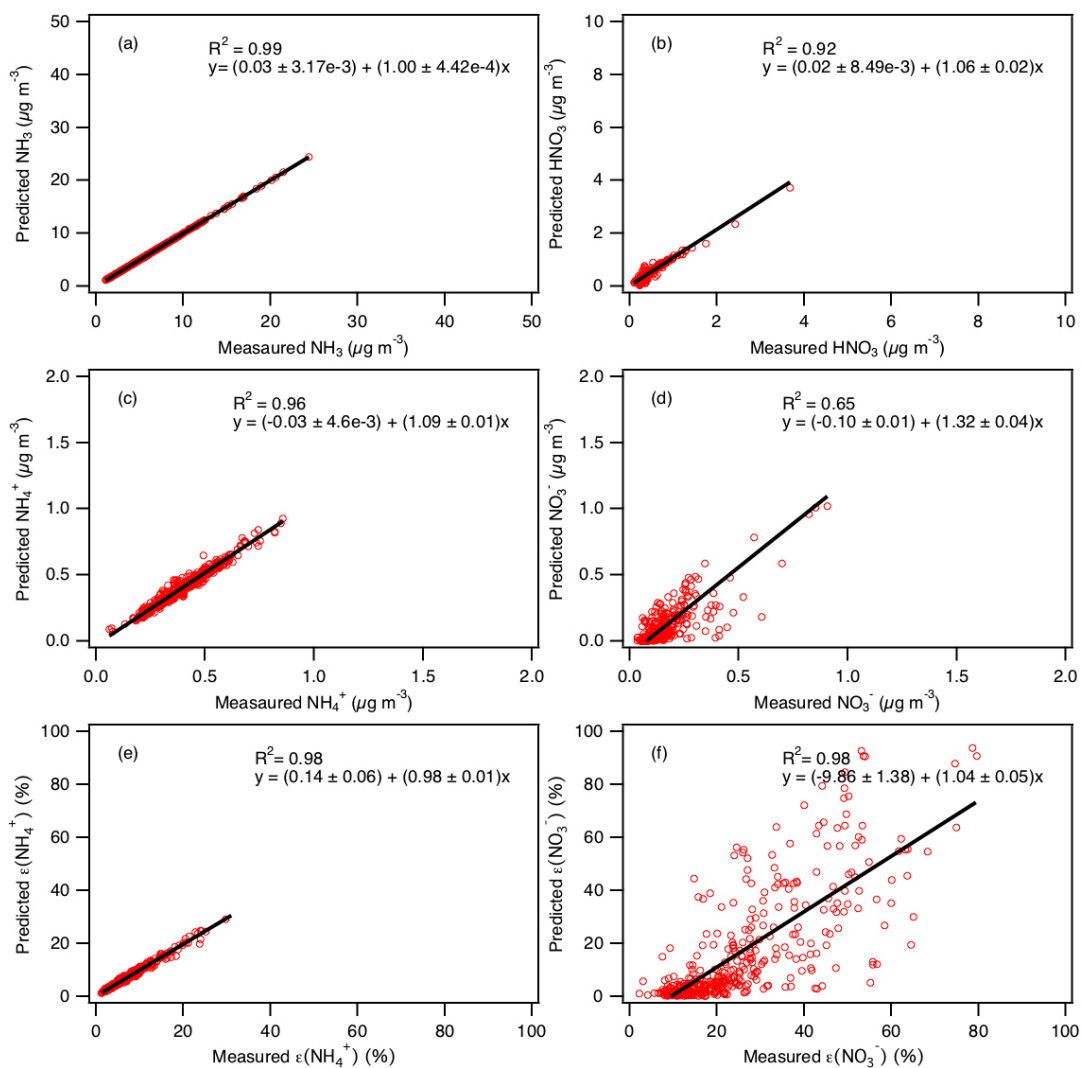
49 **Figure S6:** Aerosol (panels a to d)  $\text{SO}_4^{2-}$ , (panels e to h)  $\text{NO}_3^-$ , and (i)  $\text{NH}_4^+$  comparisons between  
50 AMS, PILS-IC, PILS-HPIC and filters for the entire field study. For comparisons between the  
51 AMS, PILS-IC and PILS-HPIC data (panels c, d, g and h), the measurements are averaged over 1  
52 hour intervals. For comparisons with filter data (panels a, b, e, f and i), the AMS, PILS-IC and  
53 PILS-HPIC data are averaged over 24 hour intervals. Orthogonal regression fits are shown.  
54  
55 Uncertainties in the fits are 1 standard deviation.



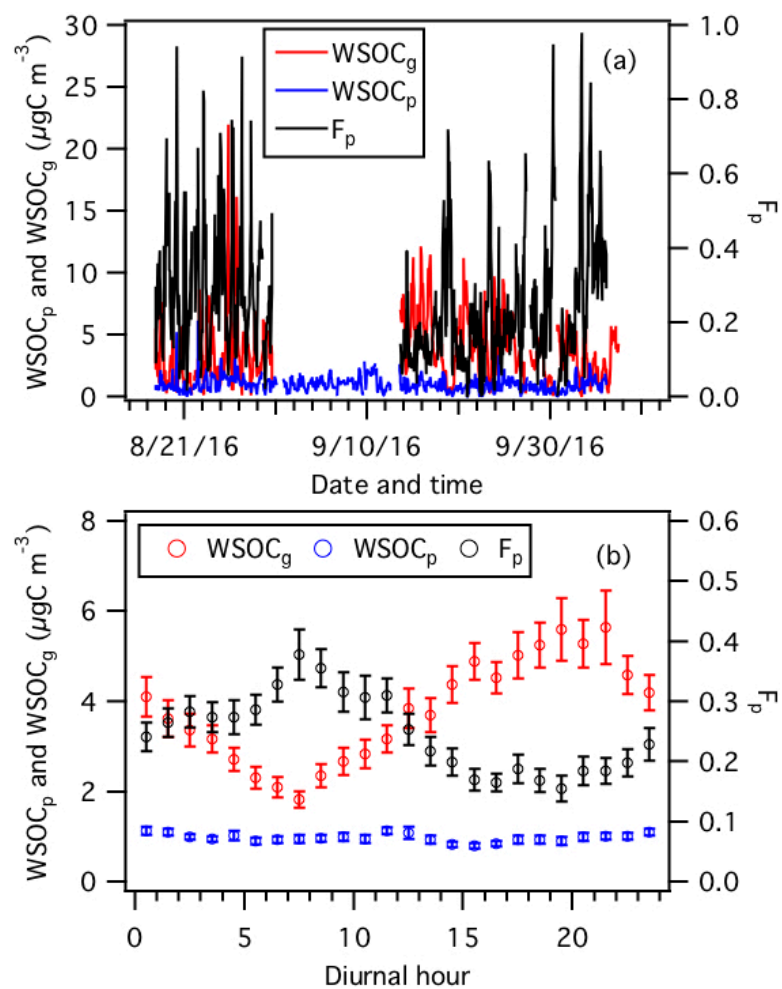
56

57 **Figure S7:** Comparison of predicted  $\text{PM}_{10}$  pH values determined using  $\text{NH}_3$ -CIMS and SEARCH  
 58 network's  $\text{NH}_3$  measurements as ISORROPIA-II model inputs. The other model inputs are the  
 59 same. The linear fit is obtained by orthogonal distance regression.



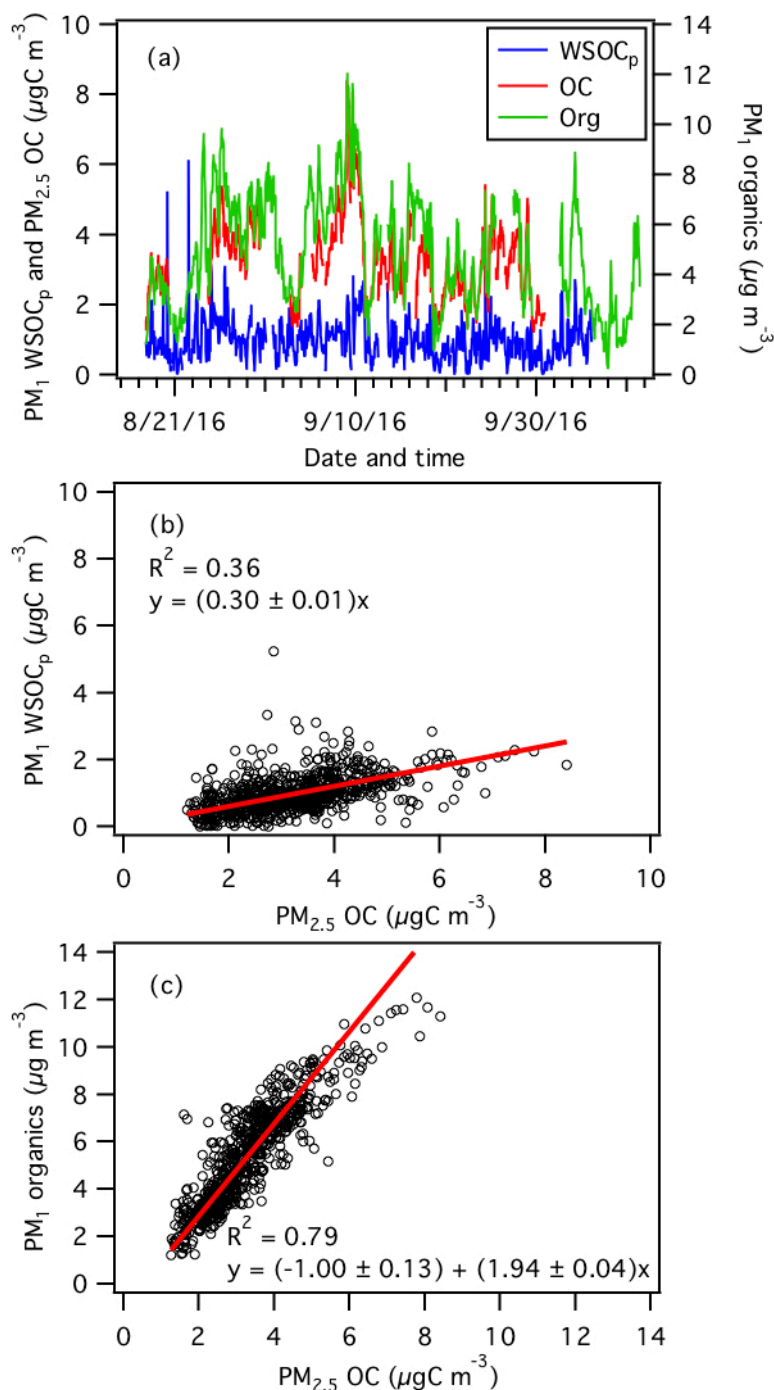


**Figure S8:** Comparisons of predicted and measured (a)  $\text{NH}_3$ , (b)  $\text{HNO}_3$ , (c)  $\text{NH}_4^+$ , (d)  $\text{NO}_3^-$ , (e)  $\epsilon(\text{NH}_4^+)$ , and (f)  $\epsilon(\text{NO}_3^-)$ . Orthogonal regression fits are shown. Uncertainties in the fits are 1 standard deviation.



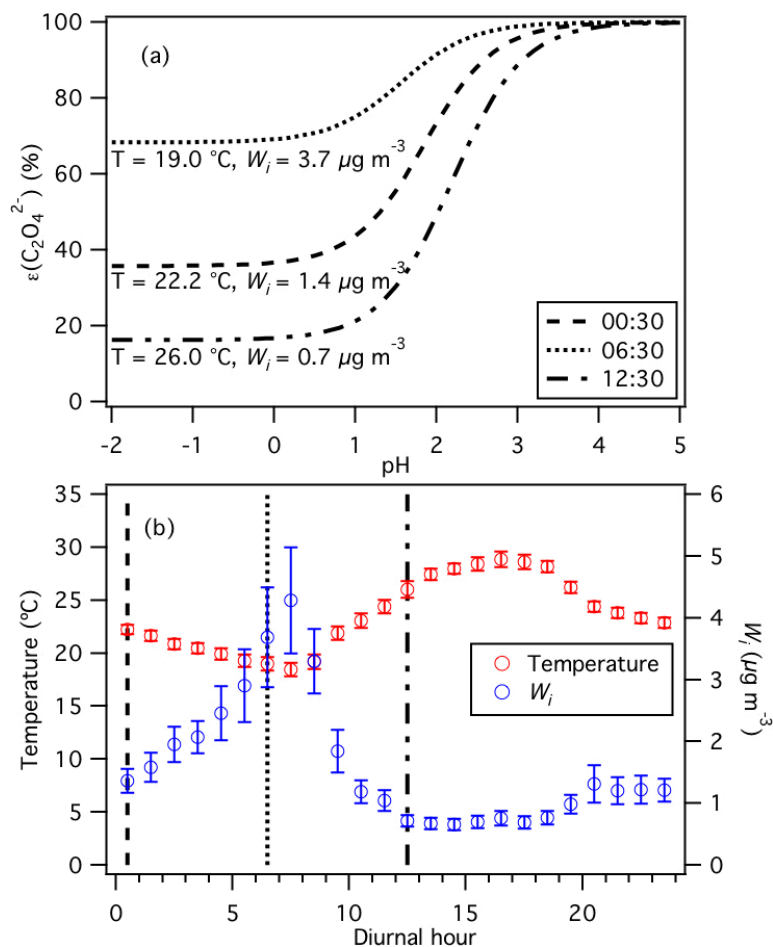
64

65 **Figure S9:** (a) Time series and (b) diurnal profiles of WSOC<sub>g</sub>, WSOC<sub>p</sub> and  $F_p$ . Dates and times  
 66 displayed are local time. All the data shown here represent averages in 1-hour intervals. Error bars  
 67 shown in panel (b) are the standard errors.  $F_p = \text{WSOC}_p / (\text{WSOC}_p + \text{WSOC}_g)$ .



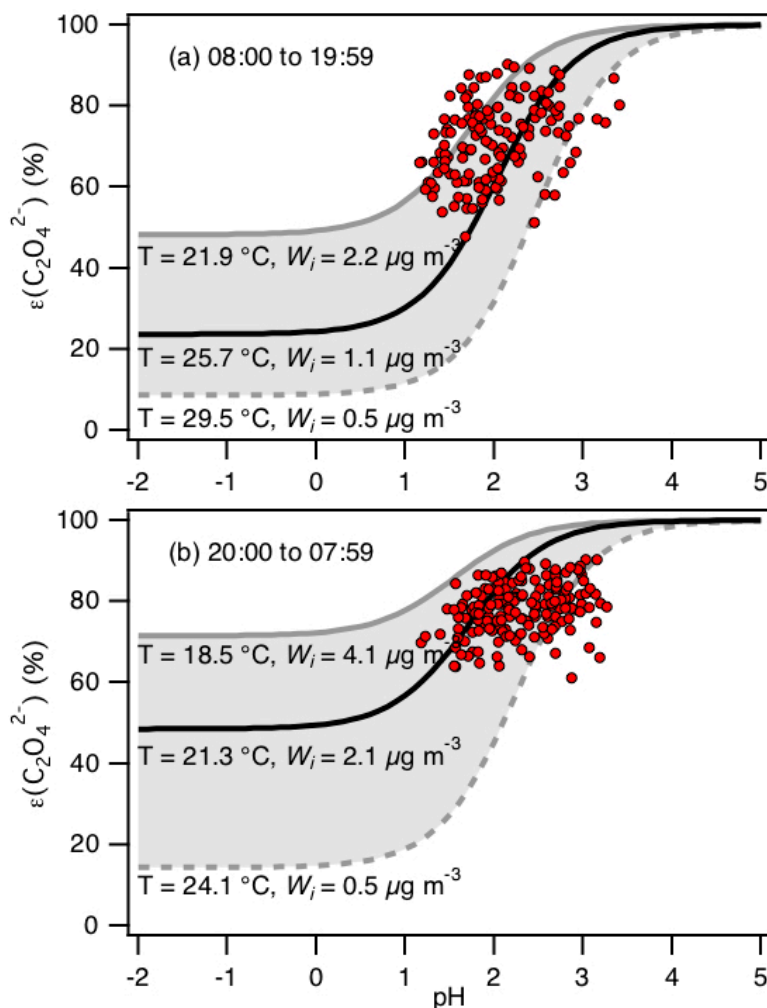
68

69 **Figure S10:** (a) Time series of AMS organics, WSOC<sub>p</sub> and OC. (b) Linear regression correlation  
70 between WSOC<sub>p</sub> and OC. (c) Linear regression correlation between AMS organics and OC. All  
71 the data shown here represent averages in 1-hour intervals. Note that OC measurements are  $PM_{2.5}$ ,  
72 while WSOC<sub>p</sub> and AMS organics measurements are  $PM_1$ . Linear fits are obtained by orthogonal  
73 distance regression.



74

75 **Figure S11:** (a) Analytically calculated S curves of  $\epsilon(\text{C}_2\text{O}_4^{2-})$  at different times of the day: 00:30,  
76 06:30 and 12:30. These S curves are calculated using values obtained from (b) the diurnal profiles  
77 of temperature and  $W_i$ . The set of 1-hour average temperatures and  $W_i$  at diurnal hours 00:30,  
78 06:30 and 12:30 is used to calculate each S curve shown in panel (a). Similar to Fig. 7, we used  
79  $\gamma_{\text{C}_2\text{H}_2\text{O}_4} = 0.0492$  (AIOMFAC predicted) and assumed that  $\gamma_{\text{H}^+ - \text{NO}_3^-} = \sqrt{\gamma_{\text{H}^+} \gamma_{\text{NO}_3^-}} = \sqrt{\gamma_{\text{H}^+} \gamma_{\text{C}_2\text{HO}_4^-}}$   
80  $= 0.265$  (ISORROPIA-II predicted) to generate these S curves.



81

82 **Figure S12:** Analytically calculated S curve of  $\varepsilon(\text{C}_2\text{O}_4^{2-})$  and ambient data from 13 September to  
83 6 October 2016 plotted against ISORROPIA-predicted particle pH. For the ambient data, a narrow  
84 range in  $W_i$  ( $0.5$  to  $4 \mu\text{g m}^{-3}$ ) and RH ( $20$  to  $90\%$ ) is chosen to be close to the analytically calculated  
85 outputs. We divided the ambient data into two sets: panel (a) 08:00 to 19:59, and panel (b) 20:00  
86 to 07:59. For both analytically calculated S curves, we used  $\gamma_{\text{C}_2\text{H}_2\text{O}_4} = 0.0492$  (AIOMFAC  
87 predicted). We also assumed that  $\gamma_{\text{H}^+}\gamma_{\text{C}_2\text{HO}_4^-} = \gamma_{\text{H}^+}\gamma_{\text{NO}_3^-}$ , and used the ISORROPIA-predicted  
88  $\gamma_{\text{H}^+}\gamma_{\text{NO}_3^-} = \sqrt{\gamma_{\text{H}^+}\gamma_{\text{NO}_3^-}} = 0.265$ . In panel (a), we used the average temperature and  $W_i$  ( $25.7 \pm 3.8$   
89  $^\circ\text{C}$  and  $1.1 \pm 1.1 \mu\text{g m}^{-3}$ ) for the data between 08:00 to 19:59 to calculate the S curve (black line).  
90 In panel (b), we used the average temperature and  $W_i$  ( $21.3 \pm 2.8^\circ\text{C}$  and  $2.1 \pm 2.0 \mu\text{g m}^{-3}$ ) for the  
91 data between 20:00 to 07:59 to calculate the S curve (black line). Grey lines in both panels are S  
92 curves calculated using one standard deviation from the average temperature and  $W_i$  for the two

datasets. In panel (a), the dotted grey line is the S curve calculated using 29.5 °C and 0.5  $\mu\text{g m}^{-3}$  while the solid grey line is the S curve calculated using 21.9 °C and 2.2  $\mu\text{g m}^{-3}$ . In panel (b), the dotted grey line is the S curve calculated using 24.1 °C and 0.5  $\mu\text{g m}^{-3}$  while the solid grey line is the S curve calculated using 18.5 °C and 4.1  $\mu\text{g m}^{-3}$ .

**Table S1:** List of gas-phase acids measured by SF<sub>6</sub>-CIMS, and their measurement uncertainties and detection limits.

Acid	Measurement uncertainty (%)	Detection limits (ppb) <sup>a</sup>
Nitric acid	13	0.20
Formic acid	12	0.03
Acetic acid	12	0.06
Oxalic acid	14	$1 \times 10^{-3}$
Butyric acid	14	0.03
Glycolic acid	22	$2 \times 10^{-3}$
Propionic acid	14	$6 \times 10^{-3}$
Valeric acid	22	0.01
Malonic acid	25	$7 \times 10^{-4}$
Succinic acid	25	$3 \times 10^{-3}$

<sup>a</sup>Detection limits are approximated from 3 times the standard deviation values ( $3\sigma$ ) of the ion signals measured during background mode. Shown here are the average detection limits of the organic acids for 2.5 min integration periods which corresponds to the length of a background measurement at a 0.04 s duty cycle for each mass.

## **S1. SF<sub>6</sub>-CIMS calibration of gas-phase HNO<sub>3</sub> and organic acids**

Detailed descriptions of post-field laboratory calibrations of HNO<sub>3</sub>, oxalic, butyric, glycolic, propionic, valeric, malonic and succinic acids can be found in Nah et al. (2018). The response of the CIMS acid signals were measured relative to the sensitivity of <sup>34</sup>SO<sub>2</sub> in these calibration measurements.

The HNO<sub>3</sub> calibration source was a permeation tube (KIN-TEK) whose emission rate was measured using UV optical absorption (Neuman et al., 2003). Solid or liquid samples of oxalic (Sigma Aldrich, ≥ 99 %), butyric (Sigma Aldrich, ≥ 99 %), glycolic (Sigma Aldrich, 99 %), propionic (Sigma Aldrich, ≥ 99.5 %), valeric (Sigma Aldrich, ≥ 99 %), malonic (Sigma Aldrich, ≥ 99.5 %) and succinic (Sigma Aldrich, 99 %) acids were used in calibration measurements. The acid sample was placed in a glass impinger, which was immersed in a water bath at a fixed temperature to provide a constant vapor pressure. For oxalic, butyric, glycolic, propionic and valeric acids, the water bath temperature was set to 0 °C. For malonic and succinic acids, the water bath temperature was set to 40 °C in order to generate large enough gas phase concentrations for calibration. 6 to 10 mL min<sup>-1</sup> of nitrogen gas (N<sub>2</sub>) was passed over the organic acid in the glass impinger. This organic acid air stream was diluted with different N<sub>2</sub> flows (1 to 5 L min<sup>-1</sup>) to obtain different mixing ratios of the organic acid. We calculated the mixing ratios based on the acid's emission rate from the impinger or the acid's vapor pressure. Emission rates of gas-phase oxalic, malonic and succinic acids from the impinger were measured by scrubbing the output of the impinger in deionized water, followed by ion chromatography analysis. We measured the vapor pressures of butyric and propionic acids at 0 °C using a capacitance manometer (MKS Instruments). We estimated the vapor pressures of glycolic and valeric acids at 0 °C using their literature vapor pressures at 25 °C and enthalpies of vaporization (Daubert and Danner, 1989; Lide, 1995; Acree and Chickos, 2010).

## **S2. WSOC<sub>p</sub> and OC**

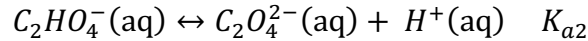
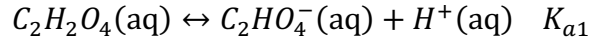
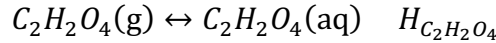
We estimated the water-soluble fraction of OC by comparing the WSOC<sub>p</sub> and OC measurements. The time series of organics, WSOC<sub>p</sub> and OC are shown in Fig. S10a. As shown in Fig. S10b, WSOC<sub>p</sub> is moderately correlated with OC at the site. The orthogonal distance regression fit suggests that 30 % of the OC is water-soluble (estimated measurement uncertainty of 19 %),



which is significantly smaller than the fraction (61 %) measured during the SOAS study (Xu et al., 2017). This difference may be due, in part, to the WSOC<sub>p</sub>/OC ratio for this study being underestimated. WSOC<sub>p</sub> are PM<sub>1</sub> measurements while OC are PM<sub>2.5</sub> measurements. This is in contrast to the SOAS study where both WSOC<sub>p</sub> and OC are PM<sub>2.5</sub> measurements. PM<sub>1</sub> organics mass concentration is highly correlated with OC and has an orthogonal distance regression slope of 1.94 (Fig. S9c), which is similar to the value (1.92) reported for the SOAS study (Xu et al., 2017).

### S3. C<sub>2</sub>H<sub>2</sub>O<sub>4</sub>- C<sub>2</sub>O<sub>4</sub><sup>2-</sup> partitioning

Here, we show the detailed derivation of equation (4) in that paper. Equilibrium between gaseous C<sub>2</sub>H<sub>2</sub>O<sub>4</sub> and particle-phase C<sub>2</sub>O<sub>4</sub><sup>2-</sup> involves the dissolution of C<sub>2</sub>H<sub>2</sub>O<sub>4</sub> into the aqueous phase (assuming particles are liquids), followed by dissociation of the dissolved C<sub>2</sub>H<sub>2</sub>O<sub>4</sub>:



for which the reaction equilibria are expressed as follows:

$$H_{C_2H_2O_4} = \gamma_{C_2H_2O_4} [C_2H_2O_4] / p_{C_2H_2O_4} \quad (1)$$

$$K_{a1} = \frac{\gamma_{H^+} [H^+] \gamma_{C_2HO_4^-} [C_2HO_4^-]}{\gamma_{C_2H_2O_4} [C_2H_2O_4]} \quad (2)$$

$$K_{a2} = \frac{\gamma_{H^+} [H^+] \gamma_{C_2O_4^{2-}} [C_2O_4^{2-}]}{\gamma_{C_2HO_4^-} [C_2HO_4^-]} \quad (3)$$

where  $H_{C_2H_2O_4}$  (mole L<sup>-1</sup> atm<sup>-1</sup>) is the Henry's law constant for oxalic acid,  $K_{a1}$  and  $K_{a2}$  (mole L<sup>-1</sup>) are the first and second acid dissociation constants for oxalic acid,  $p_{C_2H_2O_4}$  (atm) is the partial pressure of oxalic acid in the atmosphere, and  $\gamma_i$ 's are activity coefficients. In equations (1) to (3), [x] represents aqueous concentrations (mole L<sup>-1</sup>).

The total dissolved C<sub>2</sub>H<sub>2</sub>O<sub>4</sub> or particle-phase oxalate (C<sub>2</sub>O<sub>4</sub><sup>T</sup>) can be expressed as:

$$[C_2HO_4^T] = [C_2H_2O_4] + [C_2HO_4^-] + [C_2O_4^{2-}] \quad (4)$$

186 Using equations (1) to (3),  $[C_2O_4^T]$  can be expressed as:

$$187 \quad [C_2O_4^T] = H_{C_2H_2O_4} p_{C_2H_2O_4} \left( \frac{1}{\gamma_{C_2H_2O_4}} + \frac{K_{a1}}{\gamma_H + \gamma_{C_2HO_4^-} [H^+]} + \frac{K_{a1}K_{a2}}{\gamma_H + \gamma_H + \gamma_{C_2O_4^{2-}} [H^+]^2} \right) \quad (5)$$

188 The ideal gas law gives:

$$189 \quad c(C_2H_2O_4) = \frac{p_{C_2H_2O_4}}{RT} \quad (6)$$

190 where  $R$  is the gas constant,  $T$  is temperature, and  $c(x)$  represents concentration per volume of air  
191 (mole  $m^{-3}$ ). The particle-phase fraction of oxalate can then be expressed as:

$$192 \quad \varepsilon(C_2O_4^{2-}) = \frac{c(C_2O_4^T)}{c(C_2H_2O_4) + c(C_2O_4^T)} = \frac{[C_2HO_4^T]W_i}{c(C_2H_2O_4) + [C_2HO_4^T]W_i} \quad (7)$$

193 where  $W_i$  is the particle liquid water content associated with inorganic species ( $\mu g\ m^{-3}$ ; mass per  
194 volume of air). Note that the particle liquid water content associated with organic species is not  
195 considered in this case, but it can be included. Alternatively, the measured particle water can be  
196 used.

197 By putting equations (5) and (6) into equation (7),  $\varepsilon(C_2O_4^{2-})$  can be expressed as:

$$198 \quad \varepsilon(C_2O_4^{2-}) = \frac{H_{C_2H_2O_4} W_i RT \left( \frac{1}{\gamma_{C_2H_2O_4}} + \frac{K_{a1}}{\gamma_H + \gamma_{C_2HO_4^-} [H^+]} + \frac{K_{a1}K_{a2}}{\gamma_H + \gamma_H + \gamma_{C_2O_4^{2-}} [H^+]^2} \right)}{1 + H_{C_2H_2O_4} W_i RT \left( \frac{1}{\gamma_{C_2H_2O_4}} + \frac{K_{a1}}{\gamma_H + \gamma_{C_2HO_4^-} [H^+]} + \frac{K_{a1}K_{a2}}{\gamma_H + \gamma_H + \gamma_{C_2O_4^{2-}} [H^+]^2} \right)} \quad (8)$$

199 At 298 K,  $K_{a1} = 5.62 \times 10^{-2}$  mole  $L^{-1}$  and  $K_{a2} = 1.55 \times 10^{-4}$  mole  $L^{-1}$  (Haynes, 2014). Assuming  
200 that  $\gamma_{H^+} = 1$ ,  $K_{a1}K_{a2} \ll \gamma_H + \gamma_H + \gamma_{C_2O_4^{2-}} [H^+]^2$  for the conditions of our study. Hence, equation (8)  
201 can be simplified to:

$$202 \quad \varepsilon(C_2O_4^{2-}) \cong \frac{H_{C_2H_2O_4} W_i RT \left( \frac{\gamma_{H^+} \gamma_{C_2HO_4^-} [H^+] + K_{a1}}{\gamma_{C_2H_2O_4}} \right)}{\gamma_H + \gamma_{C_2HO_4^-} [H^+] + H_{C_2H_2O_4} W_i RT \left( \frac{\gamma_{H^+} \gamma_{C_2HO_4^-} [H^+] + K_{a1}}{\gamma_{C_2H_2O_4}} \right)} \quad (9)$$

203 After accounting for the SI units and substituting  $[H^+] = 10^{-pH}$ , equation (9) becomes:

$$\varepsilon(C_2O_4^{2-}) \cong \frac{H_{C_2H_2O_4} W_i RT \left( \frac{\gamma_H + \gamma_{C_2HO_4^-}}{\gamma_{C_2H_2O_4}} 10^{-pH + K_{a1}} \right) \times 0.987 \times 10^{-14}}{\gamma_H + \gamma_{C_2HO_4^-} 10^{-pH} + H_{C_2H_2O_4} W_i RT \left( \frac{\gamma_H + \gamma_{C_2HO_4^-}}{\gamma_{C_2H_2O_4}} 10^{-pH + K_{a1}} \right) \times 0.987 \times 10^{-14}} \quad (10)$$

Note that  $0.987 \times 10^{-14}$  comes from using  $R = 8.314 \text{ m}^3 \text{ Pa K}^{-1} \text{ mol}^{-1}$ , and hence needing to convert 1 atm to 1 Pa and 1 L to 1  $\mu\text{g}$ . We used the average of  $H_{C_2H_2O_4}$  values provided by Clegg et al. (1996), Compornolle and Muller (2014) and Saxena and Hildemann (1996) ( $6.11 \times 10^8 \text{ mole L}^{-1} \text{ atm}^{-1}$  at 25 °C), and accounted for the effect of temperature using equation 19 in Sander (2015). Although  $K_{a1}$  also depends on temperature, the  $K_{a1}$  value at 25 °C ( $5.62 \times 10^{-2}$ , (Haynes, 2014)) is used for all the oxalic acid S curve calculations in this paper since equations that determine temperature-dependent  $K_{a1}$  values are not available. In addition, the temperatures observed in this study are close to 25 °C.

Figure S11 provides a conceptual picture of how the relationship between  $\varepsilon(C_2O_4^{2-})$  and particle pH can change based on the time of the day. Different S curves for  $\varepsilon(C_2O_4^{2-})$  are calculated using equation (10) and 1-hour average values obtained from the diurnal profiles of temperature and  $W_i$  (specifically at 00:30, 06:30 and 12:30). The S curves are shown to differ substantially due to the diurnal variations of temperature and  $W_i$ . For example, a decrease in temperature and an increase in  $W_i$  from 00:30 to 06:30 will result in the S curve shifting to the left, which indicates that a substantially higher fraction of gas-phase oxalic acid will partition to the particle phase for a given particle pH at 06:30 compared to at 00:30. Higher  $W_i$  also increases the fraction of oxalate that partitions to the particle phase due solely to solubility, as seen from the plateau regions at low pH in Fig. S11. Conversely, an increase in temperature and a decrease in  $W_i$  from 06:30 to 12:30 will result in a considerably lower fraction of gas-phase oxalic acid partitioning to the particle phase for a given particle pH at 12:30 compared to at 06:30.

#### S4. PILS-HPIC denuder efficiency

Post-field laboratory experiments were performed to determine if disagreements between the measured and predicted molar fractions of formic and acetic acid in the particle phase were due to positive biases in particle-phase PILS-HPIC measurements as a result of less than 100 % gas removal denuder efficiency and uptake of gases in the PILS liquid system. While experiments were done solely with formic acid, similar results are expected for acetic acid. In these

experiments, liquid formic acid (Sigma Aldrich,  $\geq 99\%$ ) was diluted with ultrapure deionized water and placed in a bubbler. A flow of  $18\text{ mL min}^{-1}$  of  $\text{N}_2$  was passed through the formic acid in the bubbler. This formic acid air stream was then passed through a nafion dryer and diluted with  $52\text{ L min}^{-1}$  of  $\text{N}_2$ . Two experiments were performed. In the first experiment, the diluted formic acid air flow was introduced directly into the PILS, which was connected to a Metrohm 761 Compact IC (Metrohm AG). In the second experiment, the diluted formic acid air flow was passed through a 28 cm parallel plate carbon denuder (Sunset Lab) prior to introduction into the PILS-IC system.

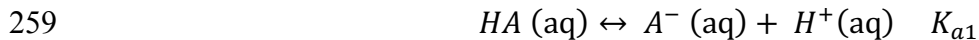
Our experiments showed that the IC detected formate when the diluted formic acid air flow was introduced directly into the PILS-IC system. IC analysis revealed that the gas-phase formic acid concentration was  $\sim 75\text{ }\mu\text{g m}^{-3}$ . However, no formate was detected above the limit of detection ( $0.02\text{ }\mu\text{g m}^{-3}$ ) when the diluted formic acid air flow was passed through the carbon denuder prior to introduction into the PILS-IC system. These measurements indicated that the carbon denuder has a  $\geq 99.97\%$  formic acid gas removal efficiency. Hence, these experiments indicate that the carbon denuder removes the formic acid gas completely. We conclude that disagreements between the measured and predicted molar fractions of formic and acetic acid in the particle phase were not due to positive biases in particle-phase formate and acetate PILS-HPIC measurements as a result of less than 100 % gas removal denuder efficiency.

## S5. Particle-phase formic and acetic acid dimers

Previous studies have shown that formic and acetic acid dimers may form in the aqueous phase (Schrier et al., 1964; Gilson et al., 1997; Chen et al., 2008). If the aforementioned acid dimers are present in aerosols, equilibrium between gas-phase formic/acetic acid (denoted as HA) and particle-phase formate/acetate (denoted as  $\text{A}^-$ ) will differ from that predicted assuming no dimers existed, as done in the main text.

The dissolution of HA into the aqueous phase (assuming particles are liquids), followed by the formation of particle-phase dimers (denoted as  $((\text{HA})_2)$  and dissociation of the dissolved HA:





for which the reaction equilibriums are expressed as follows:

$$H_{HA} = \gamma_{HA}[HA]/p_{HA} \quad (11)$$

$$K_{dim} = \frac{[(HA)_2]}{[HA]^2} \quad (12)$$

$$K_{a1} = \frac{\gamma_{A^-}[A^-]\gamma_{H^+}[H^+]}{\gamma_{HA}[HA]} \quad (13)$$

where  $H_{HA}$  (mole L<sup>-1</sup> atm<sup>-1</sup>) is the Henry's law constant for formic or acetic acid,  $K_{a1}$  (mole L<sup>-1</sup>) is the first acid dissociation constants for formic or acetic acid,  $p_{HA}$  (atm) is the partial pressure of formic or acetic acid in the atmosphere,  $K_{dim}$  (L mole<sup>-1</sup>) is the dimerization constant, and  $\gamma_i$ 's are activity coefficients. In equations (11) to (13), [x] represents aqueous concentrations (mole L<sup>-1</sup>).

The total dissolved formate or acetate ( $A^T$ ) can be expressed as:

$$[A^T] = [HA] + [A^-] + [(HA)_2] \quad (14)$$

Using equations (11) to (13),  $[A^T]$  can be expressed as:

$$[A^T] = H_{HA}p_{HA} \left( \frac{1}{\gamma_{HA}} + \frac{K_{a1}}{\gamma_{H^+}\gamma_{A^-}[H^+]} + \frac{K_{dim}H_{HA}p_{HA}}{\gamma_{HA}\gamma_{HA}} \right) \quad (15)$$

The ideal gas law gives:

$$c(HA) = \frac{p_{HA}}{RT} \quad (16)$$

where  $R$  is the gas constant,  $T$  is temperature, and  $c(x)$  represents concentration per volume of air (mole m<sup>-3</sup>). The particle-phase fraction of formate or acetate can then be expressed as:

$$\varepsilon(A^-) = \frac{c(A^T)}{c(HA) + c(A^T)} = \frac{[A^T]W_i}{c(HA) + [A^T]W_i} \quad (17)$$

where  $W_i$  is the particle liquid water content associated with inorganic species (μg m<sup>-3</sup>; mass per volume of air). Particle liquid water content associated with organic species is not considered in this case, but it can be included. Alternatively, the measured particle water can be used.

By putting equations (15) and (16) into equation (17), and accounting for the SI units,  $\varepsilon(A^-)$  can ultimately be expressed as:

$$\varepsilon(A^-) = \frac{H_{HA} W_{iRT} \left( \frac{1}{\gamma_{HA}} + \frac{K_{a1}}{\gamma_H + \gamma_{A^-} 10^{-pH}} + \frac{K_{dim} H_{HA} p_{HA}}{\gamma_{HA} \gamma_{HA}} \right) \times 0.987 \times 10^{-14}}{1 + H_{HA} W_{iRT} \left( \frac{1}{\gamma_{HA}} + \frac{K_{a1}}{\gamma_H + \gamma_{A^-} 10^{-pH}} + \frac{K_{dim} H_{HA} p_{HA}}{\gamma_{HA} \gamma_{HA}} \right) \times 0.987 \times 10^{-14}} \quad (18)$$

At 298 K,  $K_{a1}$  values are  $1.78 \times 10^{-4}$  mole L<sup>-1</sup> and  $1.75 \times 10^{-5}$  mole L<sup>-1</sup> for formic and acetic acid, respectively (Haynes, 2014).  $K_{dim}$  values are 0.56 L mole<sup>-1</sup> and 0.92 L mole<sup>-1</sup> for formic and acetic acid, respectively (Chen et al., 2008). Temperature-dependent  $H_{HA}$  values for formic and acetic acid can be obtained from Sander (2015).  $p_{HA}$  can be calculated from the measured gas-phase formic or acetic acid concentrations ( $\mu\text{g m}^{-3}$ ) and the ideal gas law. We used the web version of AIOMFAC ([www.aiomfac.caltech.edu](http://www.aiomfac.caltech.edu)) (Zuend et al., 2008; Zuend et al., 2011; Zuend et al., 2012) to compute study-averaged  $\gamma_{HA}$  values of 0.334 and 2.150 for formic and acetic acid, respectively. We assumed that  $\gamma_H + \gamma_{A^-} = \gamma_H + \gamma_{NO_3^-} = 0.07$  for both formic and acetic acid.

Comparison of S curves generated from equation (18) and those generated from equations (5) and (6) in the main text (which assumed that no dimers existed) showed that accounting for the presence of acid dimers increased predicted  $\varepsilon(A^-)$  values by less than 1 % for particle pH 0.9 to 3.8 (i.e., pH values in this study). S curves generated by equation (18) also do not match our measured molar fractions of formic and acetic acid in the particle phase. This analysis shows that the molar fractions of formic and acetic acid in the particle phase do not change substantially when the presence of particle-phase acid dimers is accounted for due to the somewhat low  $H_{HA}$  values for formic and acetic acid. Hence, disagreements between the measured and predicted molar fractions of formic and acetic acid in the particle phase are not due to the presence of particle-phase formic and acetic acid dimers.

305   **References**

- 306   Acree, W., and Chickos, J. S.: Phase Transition Enthalpy Measurements of Organic and  
307   Organometallic Compounds. Sublimation, Vaporization and Fusion Enthalpies From 1880 to  
308   2010, *J. Phys. Chem. Ref. Data*, 39, 942, 10.1063/1.3309507, 2010.
- 309   Chen, J. H., Brooks, C. L., and Scheraga, H. A.: Revisiting the carboxylic acid dimers in aqueous  
310   solution: Interplay of hydrogen bonding, hydrophobic interactions, and entropy, *Journal of*  
311   *Physical Chemistry B*, 112, 242-249, 10.1021/jp074355h, 2008.
- 312   Clegg, S. L., Brimblecombe, P., and Khan, L.: The Henry's law constant of oxalic acid and its  
313   partitioning into the atmospheric aerosol, *Idojaras*, 100, 51-68, 1996.
- 314   Compernelle, S., and Muller, J. F.: Henry's law constants of diacids and hydroxy polyacids:  
315   recommended values, *Atmos. Chem. Phys.*, 14, 2699-2712, 10.5194/acp-14-2699-2014, 2014.
- 316   Daubert, T. E., and Danner, R. P.: Physical and thermodynamic properties of pure chemicals: data  
317   compilation, Taylor & Francis, Washington, DC, 1989.
- 318   Gilson, M. K., Given, J. A., Bush, B. L., and McCammon, J. A.: The statistical-thermodynamic  
319   basis for computation of binding affinities: A critical review, *Biophysical Journal*, 72, 1047-1069,  
320   10.1016/s0006-3495(97)78756-3, 1997.
- 321   Haynes, W. M.: CRC handbook of chemistry and physics: A ready-reference book of chemical  
322   and physical data. , Boca Raton: CRC Press, 2014.
- 323   Lide, D. R.: CRC handbook of chemistry and physics: a ready-reference book of chemical and  
324   physical data, CRC Press, Boca Raton, FL, 1995.
- 325   Nah, T., Ji, Y., Tanner, D. J., Guo, H., Sullivan, A. P., Ng, N. L., Weber, R. J., and Huey, L. G.:  
326   Real-time measurements of gas-phase organic acids using SF<sub>6</sub>- chemical ionization mass  
327   spectrometry, *Atmos. Meas. Tech. Discuss.*, 2018, 1-40, 10.5194/amt-2018-46, 2018.
- 328   Neuman, J. A., Ryerson, T. B., Huey, L. G., Jakoubek, R., Nowak, J. B., Simons, C., and  
329   Fehsenfeld, F. C.: Calibration and evaluation of nitric acid and ammonia permeation tubes by UV

330 optical absorption, *Environmental Science & Technology*, 37, 2975-2981, 10.1021/es0264221,  
 331 2003.

332 Sander, R.: Compilation of Henry's law constants (version 4.0) for water as solvent, *Atmos. Chem.*  
 333 *Phys.*, 15, 4399-4981, 10.5194/acp-15-4399-2015, 2015.

334 Saxena, P., and Hildemann, L. M.: Water-soluble organics in atmospheric particles: A critical  
 335 review of the literature and application of thermodynamics to identify candidate compounds,  
 336 *Journal of Atmospheric Chemistry*, 24, 57-109, 10.1007/bf00053823, 1996.

337 Schrier, E. E., Pottle, M., and Scheraga, H. A.: The Influence of Hydrogen and Hydrophobic Bonds  
 338 on the Stability of the Carboxylic Acid Dimers in Aqueous Solution, *Journal of the American*  
 339 *Chemical Society*, 86, 3444-3449, 10.1021/ja01071a009, 1964.

340 Xu, L., Guo, H. Y., Weber, R. J., and Ng, N. L.: Chemical Characterization of Water-Soluble  
 341 Organic Aerosol in Contrasting Rural and Urban Environments in the Southeastern United States,  
 342 *Environmental Science & Technology*, 51, 78-88, 10.1021/acs.est.6b05002, 2017.

343 Zuend, A., Marcolli, C., Luo, B. P., and Peter, T.: A thermodynamic model of mixed organic-  
 344 inorganic aerosols to predict activity coefficients, *Atmos. Chem. Phys.*, 8, 4559-4593,  
 345 10.5194/acp-8-4559-2008, 2008.

346 Zuend, A., Marcolli, C., Booth, A. M., Lienhard, D. M., Soonsin, V., Krieger, U. K., Topping, D.  
 347 O., McFiggans, G., Peter, T., and Seinfeld, J. H.: New and extended parameterization of the  
 348 thermodynamic model AIOMFAC: calculation of activity coefficients for organic-inorganic  
 349 mixtures containing carboxyl, hydroxyl, carbonyl, ether, ester, alkenyl, alkyl, and aromatic  
 350 functional groups, *Atmos. Chem. Phys.*, 11, 9155-9206, 10.5194/acp-11-9155-2011, 2011.

351 Zuend, A., Marcolli, C., Luo, B. P., and Peter, T.: A thermodynamic model of mixed organic-  
 352 inorganic aerosols to predict activity coefficients (vol 8, pg 4559, 2008), *Atmos. Chem. Phys.*, 12,  
 353 10075-10075, 10.5194/acp-12-10075-2012, 2012.

354